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Title: Impacts of long-range transported mineral dust on summertime convective cloud and precipitation: a case study over the Taiwan region

Authors: Yanda Zhang, Fangqun Yu, Gan Luo, Jiwen Fan, and Shuai Liu.

We sincerely thank the referee for the detailed reviews and constructive comments which help to improve the manuscript. Below we respond to the comments in detail (*Referee's comments are in Italic*), and the revised manuscript with track-changes is attached after the replies to RCs.

Response to referee comment #1.

Review of "Impacts of long-range transported mineral dust on summertime convective cloud and precipitation: a case study over the Taiwan region" by Yanda Zhang et al.

Recommendation: Minor Revisions

The manuscript "Impacts of long-range transported mineral dust on summertime convective cloud and precipitation: a case study over the Taiwan region" mainly studies a severe precipitation event impacted by dust over Taiwan region in 2006 by using the WRF with properly cloud microphysics parameterizations and immersion freezing parameterization. In general, the paper is well written and presented in a logical way. It is a timely and important piece of work, and of general interest for cloud-aerosol interaction, extreme precipitation and so on. I therefore recommend publication of this paper in Atmospheric Chemistry and Physics after minor revisions. My comments are listed as follows:

Many thanks to the reviewer for the comments and suggestions. Please find the point-by-point responses below, and the changes to the manuscript are given by the line numbers of the revised draft with track changes.

Comments:

1. In Figure 1a: Dust are usually in coarse mode, and MERRA-2 can provide dust aerosol loadings. Why did authors just use PM_{2.5}? I think it cannot demonstrate the aerosol type.

Thanks for the comment.

In this study, we use the MERRA-2 reanalysis of dust mass mixing ratio (inst3_3d_aer_Nv) which provides the mass concentration of the dust aerosol in five size bins.

Previous studies suggest that the mass and number concentrations of dust aerosol are generally controlled by particles in coarse and fine modes, respectively (Hoffmann et al., 2008; Kaaden et al., 2009; Mahowald et al., 2014; Denjean et al., 2016). Since this study focuses more on the dust number concentration, we chose the first two bins (diameters 1.46 and 2.8 μm) of the MERRA-2 dust mass mixing ratio to qualitatively compare with the dust mass simulation (diameter from 0.5 to 2.5 μm) from GEOS-Chem-APM.

The Data section is modified to clarify this point (lines 105-110), and adjustments are made accordingly. Related references are added.

2. It should be better to add Morr2 and Morr2-ec runs in Table 1.

Thanks for the suggestion. The comparisons of Morr2 and Morr2-ec runs are added as Table 2.

- 3. There are two peaks of rain rate in this precipitation event. Large part of dust might be rainout during P1, so dust may not invigorate the convection during P2. Authors should consider about that how long the initial dust could affect development of subsequent clouds and precipitation.*

This is a good point. Following the study of Fan et al. (2014), the WRF model is used to study the dust effects, and the dust loss by convective rainout is not considered in this study (as described in Section 2.2.1).

The rainout may impact the dust loading and influence the dust-cloud interactions. However, the convective precipitation in the first case (P1) is a micro-mesoscale system, the long-range transport of dust at a large scale may contribute to dust loading as a supply. Thus we think the dust aerosol should still be able to influence the convection during P2.

In our further study on dust and dust-cloud interactions, the WRF-Chem model will be used, the rainout and dry deposition will be considered as the sink of atmospheric dust.

- 4. In Line219: “about 24% stronger than Morr2-Org and Clean runs”, is that “24%” for both runs?*

Yes, the “24%” is for both Morr2-Org and Clean runs. During P1, the maximum rain rates simulated by Morr2-Org and Clean runs occur at different times but with similar intensities.

- 5. The aerosol invigoration effect is argued to come from the enhanced latent heating when large amounts of liquid water freeze after being transported above the 0C level by convective updrafts. It would be much clearer if author mark freezing levels on the latent heating rate profiles in Figure 9 and 11.*

Thanks for the suggestion. The heights of 0°C (freezing level) and -38°C (homogeneous freezing level) are added in Figs. 7, 9, and 11.

- 6. In Figure 9(b): Why latent heating rate in clean condensation are larger than that in dusty condensation? And is there any mistake in this caption?*

We think that this weaker condensation latent heating rate between ~ 5–7 km over the strong convective grids in the SBM-Dusty case could be caused by the dryer updrafts as a result of stronger condensation below the layer (< 5 km) consuming more water vapor (Fig. S3a). The dryer updraft reduces the vapor mixing ratio at ~5–7 km, leading to weaker condensation and latent heating at this level than in the clean condition (Fig. 9b).

This weaker latent heating rate in the dusty condition caused by the dryer updraft is limited within the strong convective grids. Fig. S3b shows that, over the whole High-Pcp area, the dust effect leads to the enhanced latent heating at all altitudes, consistent with the enhanced cloud hydrometers in the dusty condition (Fig. 7).

Fig. S3 give below is added to the Supplement, and the related analysis is added to the manuscript (lines 320–330). The mistakes in the caption of Fig. 9 are revised.

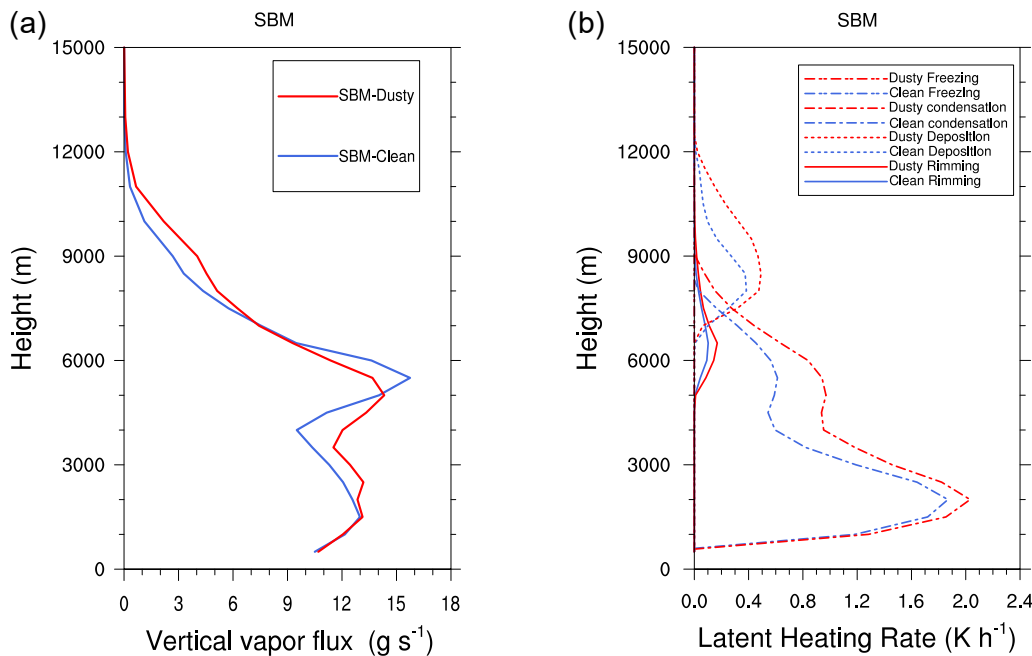


Figure S3. In the High-Pcp area during P2: (a) the vertical vapor flux averaged over grids with the top 10 percentiles of the updrafts (90th to 100th) simulated by the SBM-Clean (blue line) & Dusty (red line) runs; (b) vertical profiles of latent heating rates averaged among precipitation grids simulated by the SBM-Clean (blue lines) & Dusty (red lines) runs.

7. When talking about long range dust transportation and dust-cloud interactions over East Asia, the authors could cite following references, e.g.

- 1) Li Z., Y. Wang, J. Guo, et al. 2019: East Asian study of tropospheric aerosols and their impact on regional clouds, precipitation, and climate (EASTAIR(CPC)). *Journal of Geophysical Research: Atmospheres*. 124 (23), 13026-13054. DOI: 10.1029/2019JD030758.
- 2) Wang W., J. Huang, P. Minnis, et al. 2010: Dusty cloud properties and radiative forcing over dust source and downwind regions derived from A-Train data during the Pacific Dust Experiment. *Journal of Geophysical Research: Atmospheres*. 115. DOI:10.1029/2010JD014109.
- 3) Fu Q., T. Thorsen, J. Su, et al. 2009: Test of Mie-based single-scattering properties of non-spherical dust aerosols in radiative flux calculations. *Journal of Quantitative Spectroscopy & Radiative Transfer*. 110 (14-16), 1640-1653. DOI:10.1016/j.jqsrt.2009.03.010.

Sincere thanks! The recommended papers are now cited in the manuscript (lines 30-50).

Response to referee comment #2

In this study, the authors investigate the impact of dust aerosols on convective clouds in summertime over the mountain ranges in Taiwan. The selected case with heavy rainfall occurred in July 2006. They used the global WRF model with two different microphysics schemes, the spectral-bin (SBM) and the bulk Morrison (Morr2) scheme. They performed model simulations with and without a coupling of dust to the immersion freezing process using both microphysics schemes to analyze dust effects on convective clouds. The Morr2 scheme uses a saturation adjustment approach, which is replaced in one model simulation by explicit calculations of evaporation and condensation of cloud droplets. The model results are evaluated with observations and the MERRA-2 reanalysis. I thank the authors for this nice article and recommend it for publication after minor revisions. Please find some general minor comments below and some detailed comments and suggested corrections in the commented manuscript.

Sincere thanks to the reviewer for the comments and suggestions. Please find the point-by-point responses below, and the changes to the manuscript are given by the line numbers in the revised draft with track changes.

Comments:

8. Introduction: You should mention somewhere the typical lifetime of dust particles.

Thanks for the comment. The dust aerosol lifetime related content and references are added in the Introduction and related papers are cited (lines 25–26).

9. L45: You mentioned that dust in the mountain ranges of Taiwan originates from different source regions. You could mention that the chemical composition of dust differs for different source regions.

Thanks for the suggestion. Previous studies use multiple methods to study the sources of dust over the Taiwan region, including backward trajectory analyses, satellite observation, site measurement, model simulation, and, as mentioned, the chemical composition of dust. The related description is added (lines 56–58)

10. L65: Reference for long-range transport needed.

As introduced in this paper, previous studies indicate that the atmospheric dust over Taiwan is dominated by the long-range transport of dust and the local dust source is neglectable in most cases. Also, there is no local dust emission over the Taiwan area in the GEOS-Chem model. Thus, we think the dust event on July 8, 2006, is a long-range transport case. We cannot find related references for long-range transport in this specific case. We have slightly modified the text to avoid confusion.

11. L76: N_d is usually used for the cloud droplet number concentration and it is a bit confusing to use this here for the dust concentration.

Thanks for the comments. As this study is closely related to our previous study of the analysis of dust-cloud-precipitation interactions over the Taiwan region (Zhang et al., 2020), we use “ N_d ” to represent the dust number concentration to keep the two papers consistent. We will use other abbreviations to avoid confusing in our future work.

12. L78: You mean the N_d output are daily and hourly means, right?

Yes. The daily mean N_d is calculated by averaging the hourly model output, and related adjustments are made (line 87).

13. L82: Are there any conclusions about the model performance of the GEOS-Chem model which you can draw from previous studies?

Thanks for the suggestion. In our previous studies on the dust properties over the Taiwan region (Zhang et al., 2019; Zhang et al., 2020a), the comparison between the satellite and surface site observations indicates that the GEOS-Chem model can reasonably capture the occurrence of dust events and variations of dust concentration. The corresponding figure and text are updated in the revision (lines 93–94)

14. L137: The long-term mean is for the entire July 2006, right?

Sorry for the unclear statement in the manuscript. Here the “long-term mean” is for all the clean days (with $N_d < \text{the } 50^{\text{th}} \text{ percentile}$) in July over the last 30 years (1989–2018). An adjustment has been made to clarify this (line 155).

15. L148: Why do you consider only fine dust PM_{2.5}?

This is a good question.

Here we use the PM_{2.5} dust mass mixing ratio to quantitatively evaluate the GEOS-Chem dust simulation, this is because previous studies suggest that the mass and number concentrations of dust aerosol are generally controlled by particles in coarse and fine modes, respectively (Hoffmann et al., 2008; Kaaden et al., 2009; Mahowald et al., 2014; Denjean et al., 2016). As this study concentrates more on the dust number concentration, we chose the first two bins (diameters 1.46 and 2.8 μm) of the MERRA-2 dust mass mixing ratio to qualitatively compare with the dust mass simulation (diameter from 0.5 to 2.5 μm) from GEOS-Chem-APM.

The supplement is added in the Data section to clarify this point (lines 105–110), and adjustments are made accordingly. Related references are added.

16. L175: You mentioned the constant droplet concentration assumption in the Morr2 scheme. How about the SBM scheme?

In the SBM microphysics scheme, the lognormal modes are used to represent the concentration, mean radius, and modal width of CCN initial and boundary conditions. CCN is activated to be cloud droplets at the cloud base and in the cloud, according to the supersaturation and temperature conditions. This different treatment of cloud droplet is one of many different parameterizations between the Morr2 and SBM schemes, thus is not mentioned in the manuscript.

17. L186: Morr2-Clean and Morr2-Dusty look similar in terms of predicting a double-center in the rainfall distribution.

Thanks for the comments. In Fig.2, the simulated precipitation patterns by the Morr2-Clean and Morr2-Dusty runs do look similar. The related contents have been revised in the manuscript (lines 201–205).

18. L197: You mean the CCN activation by aerosols, since the Morr2 scheme has this constant droplet concentration, right?

Yes, the cloud droplet number concentration (CDNC) in the Morr2 is constant, and the CCN number concentration in SBM schemes is assumed and described using lognormal modes. Thus, the potential effects of CCN activation by aerosols are not considered in the Morr2 or SBM schemes in this study.

19. L200: Table 2 is missing.

Sorry for the mistake. We removed the original “Table 2” for it gave the same information as exhibited in Fig. 3. The correction is made to the draft (around line 220)

References

- Kaaden, N., Massling, A., Schladitz, A., Müller, T., Kandler, K., Schütz, L., ... & Wiedensohler, A. (2009). State of mixing, shape factor, number size distribution, and hygroscopic growth of the Saharan anthropogenic and mineral dust aerosol at Tinfou, Morocco. *Tellus B: Chemical and Physical Meteorology*, 61(1), 51-63. <https://doi.org/10.1111/j.1600-0889.2008.00388.x>
- Denjean, C., Cassola, F., Mazzino, A., Triquet, S., Chevaillier, S., Grand, N., Bourriane, T., Momboisse, G., Sellegri, K., Schwarzenbock, A., Freney, E., Mallet, M., and Formenti, P.: Size distribution and optical properties of mineral dust aerosols transported in the western Mediterranean, *Atmos. Chem. Phys.*, 16, 1081–1104, <https://doi.org/10.5194/acp-16-1081-2016>, 2016.
- Hoffmann, C., Funk, R., Sommer, M., & Li, Y. (2008). Temporal variations in PM10 and particle size distribution during Asian dust storms in Inner Mongolia. *Atmospheric Environment*, 42(36), 8422-8431. doi:10.1016/j.atmosenv.2008.08.014
- Mahowald, N., Albani, S., Kok, J. F., Engelstaeder, S., Scanza, R., Ward, D. S., & Flanner, M. G. (2014). The size distribution of desert dust aerosols and its impact on the Earth system. *Aeolian Research*, 15, 53-71, doi:10.1016/j.aeolia.2013.09.002

Impacts of long-range transported mineral dust on summertime convective cloud and precipitation: a case study over the Taiwan region

Yanda Zhang¹, Fangqun Yu¹, Gan Luo¹, Jiwen Fan², and Shuai Liu³

5 ¹Atmospheric Sciences Research Center, State University of New York, Albany, NY 12203, USA

²Atmospheric Sciences and Global Change Division, Pacific Northwest National Laboratory, Richland, WA 99354, USA

³Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing 100029, China

Correspondence to: Yanda Zhang (yzhang31@albany.edu) and Fangqun Yu (fyu@albany.edu)

Abstract. As one of the most abundant atmospheric aerosols and effective ice nuclei, mineral dust particles affect clouds and precipitation in the Earth system. Here numerical experiments are carried out to investigate the impacts of dust aerosols on summertime convective clouds and precipitation over the mountainous region in Taiwan by acting as ice-nucleating particles. We run the Weather Research and Forecasting model (WRF) with the Morrison two-moment and spectral-bin microphysics (SBM) schemes at 3 km resolution, with the using dust number concentrations from a global chemical transport model (GEOS-Chem-APM). The case study indicates that the long-range transported mineral dust, with relatively low number concentrations, can notably affect the properties of convective clouds (ice/liquid water contents, cloud top height, and cloud coverage) and precipitation (spatial pattern and intensity). The dust effects are evident during ~~the~~ strong convective periods, with significantly ~~increasing~~ increased ~~the~~ ice water contents in the mixed-phase regime via the enhanced heterogeneous freezing. With both the Morrison ~~Morr2~~ and SBM schemes, we see the invigoration effects of dust aerosols on the convective intensity through enhanced condensation and deposition latent heating. ~~In this process, the~~ The low-altitude dust particles are uplifted to the freezing level by updrafts which, in turn, enhance the convective cloud development through immersion freezing and convective invigoration. Comparing to the ~~Morr2~~ Morrison scheme, the SBM scheme predicts more realistic precipitation and different invigoration effects of dust. The differences are partially attributed to the saturation adjustment approach utilized in the bulk schemes, which leads to the ~~a~~ stronger enhancement of condensation at mid-low altitude and a weaker deposition increase at the upper level.

25 1 Introduction

~~For~~ Due to the enormous global emission rate (Zender et al., 2004; Textor et al., 2006), relatively long lifetime (days up to weeks) (IPCC, 2001; Prospero, 1999), and the long-range transport ability (Husar et al., 2004; Perry et al., 2004; Engelstaedter et al., 2006; Liu et al., 2008; Uno et al., 2009), mineral dust is one of the most abundant aerosol components in the atmosphere (Andreae et al., 1986; Carslaw et al., 2010; Kok et al., 2017). Previous studies indicate that mineral dust

30 significantly affects ~~the~~ air quality (Prospero, 1999), public health (Pope et al., 2002; Chiu et al., 2008), biogeochemical
cycles (Jickells et al., 2005), and climate systems. In the atmosphere, mineral dust is suggested to have important impacts on
the radiation budget of the Earth system through direct effects (scattering and absorbing shortwave and longwave radiations)
(Li et al., 1996; Miller and Tegen, 1998; [Fu et al., 2009](#); Zhao et al., 2011; Guo and Yin, 2015; [Li et al., 2019](#)), semi-direct
effects through changes in [the](#) atmospheric temperature structure and cloud lifetime (Hansen et al., 1997; Ackerman et al.,
35 2000; Koren et al., 2004; Huang et al., 2006).

Besides the ~~dust~~ effects on ~~the~~ radiative properties of the atmosphere, dust aerosols are also suggested to influence ~~the~~
cloud properties, enhance or suppress the clouds and precipitations by serving as cloud condensation nuclei (CCN) and ice
nuclei (IN) (Gunn and Phillips, 1957; Rosenfeld, 2000; Yin et al., 2002; Sassen et al., 2003; Chen et al., 2008; Han et al.,
2009; Min et al., 2009; Solomos et al., 2011; Seigel et al., 2013; Creamean et al., 2015; Creamean et al., 2016; Zhang et al.,
40 2020b). ~~For~~ [Due to](#) its large atmospheric loading and [good-high](#) ice-nucleating ability, mineral dust is well recognized as the
most important atmospheric ice nucleating particle (INP) (Knopf and Koop, 2006; Heymsfield et al. 2007; Klein et al., 2010;
Si et al., 2018), and a series of field observations and laboratory studies demonstrate [the dusts](#) effects on ~~the~~ ice-nucleating
and heterogeneous freezing in the cirrus and the mixed-phase clouds (Roberts and Hallett, 1968; Zimmermann et al., 2008;
Hoose and Möhler 2012; Atkinson et al., 2013; DeMott et al., 2010; DeMott et al., 2015). However, the dust effects on cloud
45 and precipitation remain to have significant uncertainties and are highly dependent on the specific cloud types and
meteorological conditions (Huang et al., 2014). Near the dust sources, the outbreaks of dust events provide high number
concentrations of INPs, leading to more ice particles at warmer temperatures through heterogeneous ice nucleation processes
(Bangert et al., 2012), which can influence the mixed-phase clouds and enhance the cloud water and precipitation (Smoydzin
et al., 2012; Min et al., 2014; Gibbons et al., 2018). Meanwhile, in the relatively dry atmosphere, water vapor competition
50 limits ice particles' growth and reduces the effective radius of cloud particles, suppressing the precipitation as a result
(Rosenfeld et al., 2001; Levin et al., 2005; Teller and Levin, 2006; Min et al., 2009; [Wang et al., 2010](#)). Over regions remote
from dust sources, ~~the~~ transport of dust ~~at-in~~ the middle and upper-troposphere ~~are-is~~ indicated to lead to dust entrainment
into ~~the~~ mixed-phase clouds, contributing to ice nuclei populations and enhancing the precipitation under the condition with
adequate water vapor inputs (DeMott et al., 2003; Muhlbauer and Lohmann, 2009; Ault et al., 2011; Creamean et al., 2013;
55 Fan et al., 2014, 2017).

In this paper, we choose the Taiwan region to study the impacts of ~~the~~ mineral dust on deep convective clouds and
precipitation. ~~Over the Taiwan area, Previous studies using backward trajectory analyses, satellite observation, site
measurement, particle chemical characteristics, and model simulation indicate that~~ the atmospheric dust loading [over the
Taiwan area](#) is dominated by the long-range transported dust from various dust sources [with different dust chemical
60 compositions](#) (East Asia, Sahara, Middle East, etc.) (Hsu et al., 2012; Lin et al., 2012; Chen et al., 2015; Chou et al., 2017;
[Wu and Huang, 2021](#)). Taiwan is also significantly affected by the summertime extreme rainfall events (Chen and Chen
2003; Xu et al., 2009; Chen et al. 2010; Lin et al., 2011; Yin et al., 2015; Zhang et al., 2020a). In our previous studies
(Zhang et al., 2020b), we found a positive correlation between the number concentration of dust particles and the

summertime rainfall intensity over Taiwan, showing that the cloud water and non-typhoon rainfall generally increase with the atmospheric dust loading over the mountain region. The main objective of this paper is to investigate the detailed physical processes underlying this positive correlation, focusing on the impacts of the low-altitude transported mineral dust (under 4 km) which has relatively low number concentrations ($< 1 \text{ cm}^{-3}$). In this study, we incorporate the DeMott et al. (2015) immersion freezing parameterization into cloud microphysics schemes in the Weather Research and Forecasting (WRF) model (Skamarock et al., 2019) to connect dust with ice nucleation, following Fan et al. (2014). We utilize this modified WRF model to simulate a heavy rainfall case over the mountain region in Taiwan, occurring on 08 July 2006, influenced by long-range transport of a dust event over the mountain region in Taiwan.

2 Data and Methodology

~~Influenced by a remote typhoon system, the circulation provided a strong water vapor transport to the mountainous region at the north western part of Taiwan island, leading to a severe precipitation event on local time (LT) 08 July 2006, with the 24 h accumulated precipitation over 220 mm in the precipitation center. At the same time, a long range transported dust event occurred over this region during the period (Fig. 1a).~~ Here we use numerical experiments and multiple datasets to study the interactions between dust aerosol and the convective clouds and precipitation. We selected a severe precipitation event over Taiwan on 8 July 2006 at local time (LT), with 24h-accumulated precipitation over 220 mm in the precipitation center. This event was influenced by a remote typhoon system, leading to a strong water vapor transport to the mountainous region at the north-western part of Taiwan Island. At the same time, a dust event occurred over this region during the period (Fig. 1a).

2.1 Data

The datasets used in this study are listed below:

(1) Dust number concentrations. In this study, the DeMott et al. (2015) immersion freezing scheme is used to describe the INP formation rate which is dependent on the atmospheric temperature and dust number concentration (diameter $> 500 \text{ nm}$) (N_d). ~~For~~ Due to the lack of measurements, N_d simulated by the GEOS-Chem model with size-resolved advanced particle microphysics (APM) (Yu and Luo, 2009) is used as the dust conditions in the WRF simulations. We run the GEOS-Chem-APM model at $2^\circ \times 2.5^\circ$ horizontal resolution with 47 vertical levels, and the daily and hourly mean N_d simulations are globally outputted at selected sites. The GEOS-Chem-APM model is driven by meteorological input from the Goddard Earth Observing System (GEOS) of the NASA Global Modeling and Assimilation Office (GMAO). In this model, dust emission is calculated with the Mineral Dust Entrainment and Deposition (DEAD) scheme (Zender et al., 2003), and dust aerosols from 30 nm to 25 μm are represented by 15 bins. The GEOS-Chem dust simulations have been widely used in studies of dust aerosol (Generoso et al., 2008; Kim et al., 2015; Ridley et al., 2016; Xu et al., 2017) and the GEOS-Chem dust simulation

over the Taiwan region is validated by our previous research, [indicating that the GEOS-Chem model can reasonably capture the occurrence of dust events and variations of dust concentrations](#) (Zhang et al., 2019; Zhang et al., 2020a).

(2) MERRA-2 dust [mass](#) mixing ratio reanalysis. The time series and vertical profile of dust aerosols can be difficult to continuously observe, especially during summertime with relatively low dust loadings. For the validation of the GEOS-Chem dust simulation, the MERRA-2 reanalysis of dust [mass](#) mixing ratio (inst3_3d_aer_Nv) (Gelaro et al., 2017; GMAO, 2015) is used as a reference for the time series and vertical distributions of the atmospheric dust aerosols. The MERRA-2 is a NASA atmospheric reanalysis for the satellite era using the ~~Goddard Earth Observing System~~ [GEOS](#) Model, Version 5 (GEOS-5) with its Atmospheric Data Assimilation System (ADAS), version 5.12.4. The MERRA project focuses on historical climate analyses for a broad range of weather and climate time scales and places the NASA EOS suite of observations in a climate context. The three-dimensional aerosol reanalysis (inst3_3d_aer_Nv) is 3 hourly data at $0.5^{\circ} \times 0.625^{\circ}$ horizontal resolution with 72 vertical layers. As the dust number concentration is not directly provided in the MERRA-2 reanalysis, the MERRA-2 dust [mass](#) mixing ratio in July 2006 is used to compare with and validate the temporal and vertical distributions of ~~GEOS-Chem APM~~ [the dust mass concentration simulated by the GEOS-Chem-APM model](#). [Previous studies suggest that the mass and number concentrations of dust aerosol are generally controlled by particles in coarse and fine modes, respectively \(Hoffmann et al., 2008; Kaaden et al., 2009; Mahowald et al., 2014; Denjean et al., 2016\). For](#) ~~Since this study concentrates~~ [this study focuses more on the dust number concentration, we chose the first two bins \(diameters 1.46 and 2.8 \$\mu\text{m}\$ \) of the MERRA-2 dust mixing ratio to qualitatively compare with the model dust mass simulation \(diameter from 0.5 to 2.5 \$\mu\text{m}\$ \).](#)

(3) Precipitation measurements. The daily precipitation data with 1-km resolution from the Taiwan Climate Change Projection Information and Adaptation Knowledge Platform (TCCIP) project is used as the rainfall observation. The TCCIP precipitation data is widely used in previous studies ~~on the~~ [for](#) extreme rainfall events in Taiwan (Chen and Chen, 2002; Chen et al., 2007; Su et al., 2012; Lin et al., 2015; Kuo et al., 2016; Henny et al., 2021). We also use the half-hourly rainfall observation at $0.1^{\circ} \times 0.1^{\circ}$ horizontal resolution from the Integrated Multi-satellite Retrievals for GPM (IMERG) (Huffman et al., 2019) as the rain rate observation.

(4) Satellite observation. In this study, the cloud measurements from the Moderate Resolution Imaging Spectroradiometer (MODIS) aboard the Aqua (EOS PM) satellites (Platnick et al., 2015) are used to compare with the model simulations. The level-2 MODIS/Aqua (MYD06_L2) cloud product consists of cloud optical and physical parameters at a spatial resolution of either 1 km or 5 km, which are derived using remotely sensed infrared (cloud top temperature, cloud top height, effective emissivity, cloud phase and cloud fraction), visible (cloud optical thickness and effective particle radius and cloud shadow effects) and near-infrared (additional information in the retrieval of cloud particle phase) solar reflected radiances.

In this study, all the data are processed into local time (LT) (Greenwich Mean Time +8).

2.2 Model configuration and experiment design

2.2.1 Immersion freezing parameterization and dust aerosol

To study the detailed physical processes of the dust-cloud-precipitation interactions in ~~this case~~ the selected case event, we conducted numerical experiments with the WRFv4.20 Model. As given in Fig.1b, two nested domains with horizontal grid spacings of 9 km (181×181) and 3 km (223×223) for Domain 1 (D01) and Domain 2 (D02), respectively, are used with 50 vertical levels up to 50 hPa. Hourly ERA5 reanalysis is used as the model initial and boundary meteorological conditions. The physical parameterization schemes used in the model include the RRTMG longwave radiation scheme, the RRTMG shortwave radiation scheme (Iacono et al., 2008), and the Mellor-Yamada-Janjic PBL scheme (Janjic 1994). The Tiedtke cumulus scheme (Tiedtke 1989) is used in D01 and turned off in D02. The simulations were initiated at 0000 UTC 07 July with 16 hours spin-up time and ended at 0000 UTC 09 July 2006. Simulations in D01 are mainly used to provide meteorological conditions for D02, and the analysis in this study is focused on the fine-resolution simulation results in D02.

In ~~the~~ summertime, the dust number concentration is relatively low over this region remote from dust sources, with the $N_d \approx 0.12 \text{ cm}^{-3}$ in the top 10% of dust events simulated by the GEOS-Chem model (Zhang et al., 2020a). During ~~this the~~ selected heavy rainfall event, the atmospheric dust loading over this region is too low to influence the optical properties of the atmosphere ($\sim 0.6 \text{ cm}^{-3}$), the direct effects of dust on solar radiation are not considered and this study focuses on the effects of dust as INPs. To investigate the impacts of dust aerosols on clouds and precipitation, we employed two cloud microphysics parameterizations: the spectral-bin microphysics scheme (SBM; Khain et al. 2004; Fan et al. 2012) and two-moment Morrison microphysics scheme (Morr2; Morrison et al., 2009). This is to evaluate how the dust effects vary with the two typical types of cloud microphysics parameterizations. Based on the SBM and Morr2 schemes, we designed six sensitivity experiments using different settings of the immersion freezing parameterizations and dust conditions (Table 1). In the control runs (Morr2-Org and SBM-Org), the default Morr2 and SBM schemes in WRF v4.2.0 with the Bigg (1953) immersion freezing parameterization are utilized without ~~consideration~~ considering of the dust effects. In the four dust-related runs (Morr2-Clean & Dusty, SBM-Clean & Dusty), to connect dust with ice nucleation, the immersion freezing parameterization of DeMott et al. (2015) is implemented to replace the original Bigg (1953) immersion freezing in both the SBM and Morr2 microphysics schemes. ~~In the Clean and Dusty runs, two vertical profiles of N_d from the GEOS-Chem APM model are used as the initial and boundary conditions of WRF simulations to represent the atmospheric dust number concentration in the clean and dusty conditions.~~ The N_d vertical profile simulated by the GEOS-Chem-APM model during this event (07–08 July 2006) is used in the Morr2-Dusty and SBM-Dusty runs, to represent the N_d condition of this long-range dust transport event (red dots in Fig. 1c); and for comparison to investigate the dust effects, the long-term mean N_d profile averaged in the clean atmosphere (with $N_d < \text{the } 50^{\text{th}} \text{ percentile of daily dust number concentrations in July, from } 1989 \text{ to } 2018$) is used in the clean runs (Morr2-Clean and SBM-Clean) (blue triangles in Fig. 1c). The vertical profiles of N_d show that, during this long-range dust transport event, the dust number concentration increased by about two orders of

magnitude from the clean atmosphere. This mineral dust transport event occurred at a low altitude (under ~3 km) with no evident impacts on N_d at the upper level (above ~6 km). Following Fan et al. (2014), the horizontal and vertical transports of dust are driven by WRF dynamic and ice formation through immersion freezing acting as the sink of dust and liquid droplets. ~~The rainout~~ Rainout and dry deposition of dust are not considered in this study. Here, only the immersion freezing is modified to depend on dust aerosol concentrations, while other ice formation processes (homogeneous, deposition freezing, etc.) remain unmodified as in the default SBM and Morr2 microphysics schemes in ~~the~~ WRF v4.2.

To validate the dust simulated by the GEOS-Chem-APM model, the MERRA-2 dust mixing ratio reanalysis is used to qualitatively compare with the dust simulation. Figure 1 a shows the time series of the GEOS-Chem-APM dust number concentration (red dashed line) and ~~PM2.5~~ dust mixing ratio (red solid line) simulations and MERRA-2 ~~PM2.5~~ dust mixing ratio reanalysis (blue line) in July 2006. The comparison shows that the time series of the GEOS-Chem dust simulation is generally consistent with the MERRA-2 reanalysis, predicting ~~the a~~ strong dust signal from 7–11 July 2006 during the selected severe precipitation event. The vertical distribution of N_d from the GEOS-Chem-APM is also consistent with that of the MERRA-2 reanalysis (Fig. S1), with the highest dust loading at ~1 km and decreasing with altitude. Overall, the GEOS-Chem-APM model captures the occurrence and vertical distribution of this ~~low-level transport of dust~~ event, and the dust profile used in the Dusty runs is reasonable to represent the actual dust conditions of this case.

2.2.2 Supersaturation-based Condensation

Previous studies indicate that the saturation adjustment approach for condensation and evaporation in the Morr2 microphysics scheme influences the simulations of the invigoration effects of aerosols on the deep convective clouds (Lebo and Seinfeld, 2011; Lebo et al., 2012; Wang et al., 2013; Khain et al., 2015; Zhang et al., 2021). To figure out the potential influences of the saturation adjustment on the dust-cloud interaction in this case, we modified the Morr2 scheme by replacing the saturation adjustment assumption with the explicit calculation of condensation and evaporation based on supersaturation. Following Lebo and Seinfeld (2011), the change of the water vapor surplus within each timestep as a result of phase changes based on the microphysical forcing is expressed as:

$$\frac{d\Delta q_v(t)}{dt} = -C\Delta q_v(t) \quad (1)$$

where $\Delta q_v = q_v - q_{v,s}$, q_v is the water vapor mixing ratio, $q_{v,s}$ is the saturation water vapor mixing ratio. C is a function of temperature (T), pressure (P), droplet mass (q_c), and number concentrations (n_c), with details given in Pruppacher and Klett (1997) and Porz et al. (2018). The supersaturation in Eq. (1) is solved by analytical integration over the time step, and the integrated water vapor surplus change is used to calculate the cloud droplet condensation. The growth equation of q_c following Porz et al. (2018) is expressed as:

$$\frac{dq_c}{dt} = \rho dn_c^{\frac{2}{3}} q_c^{\frac{1}{3}} \Delta q_v \quad (2)$$

where ρ is the air density and d is a function of pressure (P) and temperature (T) defined as:

$$d = 4\pi\left(\frac{3}{4\pi\rho_w}\right)^{\frac{1}{3}}DG \quad (3)$$

190 ρ_w is the density of liquid water, D is the diffusion term depending on T and P , and G is the term of latent heat influence.

In the Morr2-Clean and Morr2-Dusty, Eq. (2) is used to replace the saturation adjustment approach in the Morr2 scheme and integrated analytically over the time step to explicitly calculate condensation/evaporation growth of cloud droplets. These two simulations are referred to as Morr2-ec-Clean and Morr2-ec-Dusty (Table 2), respectively. Due to the constant droplet concentration assumption in the Morr2 scheme, the influences of supersaturation on the cloud condensation
 195 nuclei activation are not considered in this study.

3 Results

3.1 Simulated dust effects on precipitation

In order to investigate the influences of dust aerosols on clouds and precipitation, the simulation results of the six runs are compared and evaluated with available observations and reanalysis. Figure 2a shows the 24-h accumulated precipitation
 200 (r_{24h}) from 0000 to 2300 LT on 08 July from the TCCIP observation system. The observation shows that the rain band is located along the west coast of Taiwan, and two precipitation centers located at $\sim 23.25^\circ\text{N}$ & 120.5°E and $\sim 22.5^\circ\text{N}$ & 120.5°E , respectively, with ~~the a~~ heavy rainfall amount larger than 220 mm d^{-1} . Overall, the six model runs capture the spatial pattern of the observed precipitation on the windward slope of the mountain region, consistent with the observation. However, distributions of the rainfall centers simulated by ~~the SBM-Clean (Fig. 2c), and Morr2-Org (e), and Morr2-Clean~~
 205 ~~(f)~~ runs mismatch with the rainfall measurement. SBM-Org (b), SBM-Dusty (d), Morr2-Clean (f), and Morr2-Dusty (g) generally capture the observed double-center rainfall distribution. From Morr2-Clean to Morr2-Dusty, r_{24h} averaged over all rainfall grids ($r_{24h} > 1 \text{ mm d}^{-1}$) and the strong precipitation grids (top 10 percentiles of the r_{24h} , 90th to 100th) in the heavy precipitation area (High-Pcp) marked by the red box increased by around 10% (36 to 40 mm d^{-1}) and 15% (122 to 140 mm d^{-1}), respectively. From SBM-Clean to SBM-Dusty, r_{24h} averaged over all rainfall grids and the strong precipitation grids in the
 210 High-Pcp area increased by about 18% (44 to 52 mm d^{-1}) and 20% (128 to 153 mm d^{-1}), respectively. ~~These~~ results illustrate that both the bulk and bin schemes predict the stronger accumulated precipitations with the dust effect considered, and the SBM scheme simulates the observed precipitation pattern and intensity better compared to the Morr2 scheme. enhancement of the 24 h rainfall in better agreement with the observation. Nevertheless, the WRF model underpredicts the intensity of the rainfall centers in all six runs to differing degrees. Among them, SBM-Dusty simulates the highest rainfall amount with ~ 200
 215 mm d^{-1} in the precipitation centers, which is still weaker than the observation. Some factors could lead to this underestimation predicted by the WRF runs. One of the possible reasons is that the contributions of other natural and anthropogenic aerosols to CCN and INPS are is-not considered in the Morr2 or SBM schemes in this numerical study, and

another factor could be the relatively coarse resolution (3-km), which cannot resolve the orographically forcing and mountain-valley circulation well.

220 To quantitatively exhibit the difference among the six runs, the Taylor diagram (Fig. 3 and Table 2) is used to evaluate the simulations of the 24-h accumulated precipitation. Dusty runs simulate improved precipitation patterns compared to the Org and Clean runs, with spatial correlation coefficients (r) improved from ~ 0.56 to 0.62 and from ~ 0.55 to 0.66 with Morr2 and SBM schemes, respectively. SBM-Dusty simulates ~~the a~~ stronger precipitation more consistent with the observation, with normalized standard deviation increased from ~ 0.58 to 0.67 , and normalized bias decreased from $\sim -33\%$ to -24.7% ,
225 respectively. Although Morr2-Dusty predicts a better rainfall pattern than the Org and Clean runs, its simulation of the 24-h accumulated precipitation amount shows no apparent improvement.

As shown in Fig. 2, the strong precipitation, based on observation and simulations, is mainly located in the High-Pcp area, which is selected to investigate the precipitation and cloud properties in detail. The hourly rain rates (r_h) by the six runs and the IMERG half-hourly observation are averaged among the precipitation grids ($r_h > 0 \text{ mm h}^{-1}$). The 24-h accumulated
230 precipitation from IMERG is given in Fig. S2, with the intensity and distribution consistent with the rain gauge-radar observation from TCCIP (Fig. 2a). In Fig. 4, the observed hourly rain rate (black line) shows two peaks, the weaker one occurred in the morning (LT 06:00–10:00) (P1, maximum $r_h = 3.5 \text{ mm h}^{-1}$) and the stronger one in the afternoon (LT 11:00–15:00) (P2, maximum $r_h = 7.1 \text{ mm h}^{-1}$). The observed bimodal distribution is generally captured by all six runs. However, the simulated rain rate intensities are generally weaker than the observation, especially during the P2 period. With both ~~the~~
235 Morr2 and SBM schemes, the Dusty runs show stronger rain rate intensities at the two peak times. During the period of P1, the maximum r_h simulated by SBM-Dusty is 3.1 mm h^{-1} , about 45% stronger than SBM-Org and Clean runs; during the P2 period, SBM-Dusty simulates maximum $r_h = 5.2 \text{ mm h}^{-1}$, about 38% and 55% stronger than SBM-Org and Clean runs, respectively. SBM-Dusty also captures the observed precipitation evolution, which is ~~the a~~ weaker rainfall intensity of P1 compared with P2. During the P1 period, the maximum rain rate simulated by Morr2-Dusty is 4.2 mm h^{-1} , about 24%
240 stronger than Morr2-Org and Clean runs. During the P2 period, Morr2-Dusty simulates maximum $r_h = 4.0 \text{ mm h}^{-1}$, about 18% and 42% stronger than Morr2-Org and Clean runs. It is notable that, Morr2-Org and Clean dramatically overestimated the precipitation from LT 1600 to 2000 compared to the observation, and the false signal is reduced when the dust effect is considered (Morr2-Dusty). Although the total rainfall volumes simulated among the Morr2 runs are similar (Fig. 3), Morr2-Dusty predicts an improved precipitation development ~~compared to than~~ the Org and Clean runs with maximum rain rates
245 closer to the observation. Morr2-Dusty still simulates a stronger rainfall intensity of P1 than P2, mismatching ~~with the~~ observed precipitation evolution.

Therefore, connecting dust aerosols with ice nucleation improves the predicted intensity, distribution, and development of precipitation with both the bulk and bin schemes. In this case, the result also suggests that the SBM scheme outperforms Morr2 in simulating the rainfall intensity, distribution, and temporal evolution. The dust effects (Dusty vs. Clean) increase
250 the rainfall intensities in the heavy precipitation area and periods, and SBM simulated a stronger dust effect than Morr2.

3.2 Dust impacts on cloud properties

We further investigate the dust effects on the cloud evolution by investigating the time series of the total (ice + liquid) water content (TWC). Figure 5a–d show that WRF runs predict the time series of cloud properties with two distinct periods of convection, explaining the temporal distribution of the hourly rain rates in the observation and simulation. The comparisons between Clean and Dusty runs (Fig. 5a & c vs. 5b & d) suggest significant impacts of dust on cloud development. In the Morr2-Dusty and SBM-Dusty runs, the high N_d enhances the heterogeneous freezing process (immersion freezing), leading to the increased TWC during the strong convection periods (P1 and P2). From Morr2-Clean to Dusty runs, the maximum ice water content (IWC_{max}) increased by about 50% and 35% during P1 and P2, respectively. From SBM-Clean to Dusty runs, the maximum liquid water contents (LWC_{max}) increase by ~20% and 30%, and the IWC_{max} increase by about 70% and 60% during P1 and P2, respectively. Along with the enhanced TWC, the mean cloud top height (CTH) during P2 increases from ~12 km to 14 km with Morr2 and SBM schemes, when the long-range transport of dust is considered.

During the strong convective period (P2) with the observed precipitation peak, the MODIS Aqua satellite passed the simulation region at LT 14:00. To further investigate the dust effects on the ~~simulations of simulated~~ cloud properties, we use the MODIS observations of the cloud water path (CWP, liquid water path + ice water path) and CTH as references to compare with the Morr2 and SBM simulations in Clean and Dusty conditions. This study concentrates on the ~~mixed-phase and convective clouds~~, and the grid cells are selected with a minimum TWC of 0.05 g kg^{-1} and minimum CWP of 100 g m^{-2} (Choi et al., 2010; Huang et al., 2015). MODIS observations (Fig.6 a & f) show that, during P2, the convective cloud is distributed along the mountain region, especially over the northern part of the island. The observed high values of CWP and CTH along the south-western coast, indicate that the deep convective cloud is concentrated in the key area as given in Fig.2 (120.2°E – 120.75°E , 22.25°N – 23.25°N). Now we examine the dust effects on the convective cloud simulated with the SBM and Morr2 microphysics schemes. In comparison to the MODIS observations, the Morr2 and SBM runs generally capture the cloud distribution over the mountain region but underestimate the observed cloud coverage and cloud top height (over 15 km). This underestimation could be caused by various factors: one possible reason is ~~that~~ the limitation of the minimum TWC cannot be applied in the MODIS observation to eliminate the ~~high-altitude cirrus clouds~~; also the underprediction of CTH could be related to the relatively coarse model resolution. ~~The comparisons~~ Comparisons between Clean (Fig.6 b & d) and Dusty (Fig.6 c & e) runs show that the cloud coverages are enhanced in the key area when the dust effect is considered. At the same time, Fig.6 g–j show that the dust effect enhances the CTH from 12 km to over 14 km (Fig.6 g & i vs. Fig.6 h & j), more consistent with the MODIS measurement (CTH >15 km).

During the P2 period, the simulated rain rate, cloud coverage, and CTH are notably enhanced due to the dust effects. The change of these cloud properties can be caused by both convective invigoration and the microphysics effect (Futyan and Del Genio, 2007; Fan et al. 2013; Gibbons et al., 2018). Thus, we focus on the strong convective period P2 to further study in detail the dust effects on cloud properties.

In this numerical study, mineral dust is directly connected with the cloud ice/snow particle formation through the immersion freezing process. The changes of the hydrometeor species are analyzed to exhibit the dust microphysical effects in detail. Figure-Fig. 7 shows vertical profiles of the number concentration and mixing ratio of the hydrometeor particles averaged among the rainfall grids ($>0 \text{ mm h}^{-1}$) in the High-Pcp area in the P2 period. In Fig. 7 the Morr2 and SBM schemes simulate similar vertical distributions of cloud hydrometeors, with the largest IWC and LWC appearing at $\sim 8 \text{ km}$ and under 3 km , respectively. The Comparisons between Clean and Dusty runs (Fig. 7 blue vs. red) show notable dust effects in the mixed-phase layer ($\sim 5\text{--}11 \text{ km}$). With both the Morr2 and SBM schemes, dust aerosol remarkably enhances the formation of snow particles in the mixed-phase layer, with the snow number concentration about doubled and snow mixing ratios increased by $\sim 60\text{--}70\%$ at $\sim 8 \text{ km}$ altitude. From clean to dusty conditions, the increased in snow particles causes as the stronger riming and collision-coalescence processes, leading to the increased number and mass mixing ratios of graupel in both Morr2 and SBM runs.

In Fig.7, from Clean to Dusty conditions, the raindrop number and mass concentrations notably increase in the mixed-phase regime (around $4\text{--}9 \text{ km}$) and under the melting layer (under $\sim 5 \text{ km}$ altitude) with both Morr2 and SBM schemes. The dust aerosol not only influences the LWC but also shows impacts on the raindrop particle size distribution (PSD), which is an important factor influencing the collisional growth of raindrops and the cloud microphysics precipitation efficiency. The PSD in the SBM microphysical scheme is represented by the particular bin microphysical information, and the PSD in the Morr2 scheme is represented by gamma functions based on the predicted hydrometeor mass, number, and the fixed shape parameters (Morrison et al., 2009). The particle mass spectrum size distributions of the snow particles in the mixed-phase layer ($4\text{--}7 \text{ km}$) and raindrop under the melting layer (under 4 km) averaged in P2 are given in Fig. 8. The result shows, during the P2 period, more and larger snow particles are generated through the enhanced immersion freezing process in the dusty condition (Fig.8 a), and the melting of larger snow particle further leads to the increased raindrop size under the melting layer and enhanced surface rain rate.

The more numerous and larger hydrometeor particles generated in the dusty atmosphere could also be a result of the enhanced convective intensities due to the dust effects. As described in Section 2.2.1, N_d is only connected with the immersion freezing processes of the microphysics schemes, the direct radiative forcing by dust aerosols is not considered. The dynamic and thermodynamic effects of dust are caused through latent heating from the formation and growth of hydrometeor particles. To tease out the dust aerosol impacts on the convection, the vertical profiles of mean updraft velocities and latent heating rates averaged over the strong convective grids of the top 10th percentile updrafts (with the value greater than $\sim 0.5 \text{ m s}^{-1}$) during the P2 period are given in Fig. 9. The result shows that the dust effect leads to the an enhancement of the riming latent heating in the mixed-phase layer ($\sim 5\text{--}10 \text{ km}$) with the Morr2 and SBM schemes (Fig. 9a-b). In this process, with the dust effects considered, the stronger riming and Wegener-Bergeron-Findeisen (WBF) processes lead to increasing latent heat release in the mixed-phase regime, which enhances the convective intensities. However, the convective invigoration effects lead to different further impacts on the updrafts and condensation and deposition processes in the Morr2 and SBM runs. From Morr2-Clean to Morr2-Dusty runs, dust effects notably increase the condensation latent

heating at 4–8 km altitude (Fig. 9a), leading to the significantly increased convective intensity (by ~34%) at the same level (Fig. 9c). At the upper level (8–12 km), the Morr2 scheme predicts limited dust effects on the deposition latent heating and the convective intensity. With the SBM microphysics scheme, dust aerosol shows a weaker convective invigoration effect at the lower level comparing to that in the Morr2 runs. Meanwhile, at upper levels, the SBM scheme predicts a stronger enhancement of deposition latent heating due to the dust effect, leading to the notably increased convective intensity (increased by ~ 20% at 10 km, and about doubled at higher altitudes) (Fig. 9b & c). [Fig. 9b shows that over the intense convective grids, SBM-Dusty simulates the lower condensation latent heating at ~ 5–7 km than in the clean condition. This weaker condensation latent heating rate between ~ 5–7 km in the SBM-Dusty case could be caused by the dryer updrafts as a result of stronger condensation below the layer \(< 5 km\) consuming more water vapor \(Fig. S3a\). This weaker latent heating rate in the dusty condition caused by the updraft and the vertical vapor transport is limited within the strong convective grids. Fig. S3b shows that, over the whole High-Pcp area, the dust effect leads to the enhanced latent heating at all altitudes, consistent with the enhanced cloud hydrometers in the dusty condition \(Fig. 7\).](#)

The analysis shows the effects of dust aerosols on clouds and precipitation, enhancing the convective system via microphysical processes and latent heat release. At the same time, the convection system also has impacts on the advection and vertical transport of dust. The N_d vertical profiles (Fig. 1c) show that without the convective transport of dust, the dust loadings above the freezing level (> 6 km) are similar in both Clean and Dusty conditions ($N_d < 0.002 \text{ cm}^{-3}$). Figure 10 shows that, even with the orographic uplifting effect, the low-level transported dust aerosols are distributed below the freezing level (0 °C) without the occurrence of convections, which have no effects on warm clouds and precipitation without ice-phase processes. During the strong convection stage (P1 and P2), dust aerosols are lifted to the upper level and N_d can reach about 0.12 cm^{-3} at 9 km altitude (~ -20 °C), about two orders of magnitude higher than in the clean atmosphere, which could have considerable impacts on the cloud ice formation and enhance the development of cloud and convection in return.

The result above shows the differences in simulated dust effects on the cloud microphysical and convective properties with the Morr2 and SBM microphysics schemes. The Morr2 microphysics scheme predicts the more notable enhancement of LWC than the SBM scheme in and under the mixed-phase layer; at the upper level (~ 9–12 km), with dust effects considered, the snow and graupel mass concentrations increase more significantly with the SBM scheme than with the Morr2 scheme (Fig. 7 c vs. d). And comparing to the Morr2 scheme, the SBM scheme predicts a weaker increase of updraft at the mid-level (around 4–8 km) and a more dramatic enhancement of updraft at the upper level (~ 8–12 km) (Fig. 9). These different dust effects could be caused by the different assumptions and parameterizations of the various microphysical processes between bulk and bin microphysics schemes. Among various factors, the saturation adjustment approach in the bulk scheme is indicated to have impacts on the simulations of the aerosol-cloud interaction processes comparing to the explicit representation of the supersaturation-forced condensation growth in the bin models (Lebo and Seinfeld, 2011; Lebo et al. 2012; Khain et al., 2016; Porz et al., 2018; Zhang et al., 2021). As described in section 2.2.2, based on the Morr2-Clean and Morr2-Dusty, the saturation adjustment approach is replaced with the explicit condensation growth (as Morr2-ec-Clean and Morr2-ec-Dusty).

~~The comparison~~Comparing shows the differences in latent heating (Fig. 9a vs. Fig. 11a), updraft (Fig. 9c vs. Fig. 11b), and hydrometeor mixing ratio (Fig. 7c vs. Fig. 11c) simulated by the Morr2 and Morr2-ec runs. ~~The results~~ demonstrates the differences in the prediction of dust effects, ~~that~~The explicit supersaturation scheme produces ~~the a~~ weaker enhancement of the condensation (increased by ~30% from Morr2-ec-Clean to Morr2-ec-Dusty) than the saturation adjustment approach (increased by ~60% from Morr2-Clean to Morr2-Dusty) at the mid-level (~ 4–7 km). And at the upper level (~ 8–12 km), Morr2-ec predicts ~~the an~~ enhanced deposition latent heating from clean to dusty conditions (Fig. 11a), which is not predicted by the saturation adjustment scheme (Fig. 9a). Accordingly, from Clean to Dusty conditions, Morr2-ec predicts a smaller increase in liquid hydrometeor mixing ratio at low-altitude and a larger enhancement of the ice particle mass concentrations at the upper level (Fig. 11c). The latent heating rate can influence the buoyancy and updraft strength, thus the Morr2-ec predicts a smaller increase in updraft strength at the mid-level (~ 4–7 km) and more obviously invigorated convection at upper altitude (~ 8–12 km). The possible reason for the differences between the saturation adjustment (Morr2) and explicit supersaturation approach (Morr2-ec) is that the use of saturation adjustment approach turns all excess water vapor into liquid water at the end of each time step, leading to the stronger enhancement of condensation and latent heating at mid-low levels, and limits buoyancy increase and convective invigoration at upper levels, which is demonstrated in previous studies (Lebo et al., 2012; Grabowski and Morrison, 2017). The different simulations from the Morr2, SBM, and Morr2-ec scheme runs suggest that the use of saturation adjustment approach in the bulk scheme is one of the important factors contributing to the different dust effects predicted by the Morr2 and SBM runs. The accumulated precipitation and rain rate (Fig. [S3-S4](#) & [S4-S5](#)) from Morr2-ec runs are still different from the SBM runs, indicating that the saturation adjustment approach cannot explain all the differences in the predicted dust-cloud interactions by the bin and bulk schemes.

4 Conclusions and Discussion

This study explores the influences of ~~the~~ long-range transported (under ~5 km) mineral dust on the convective clouds and precipitation. We have conducted numerical simulations of a heavy rainfall case occurring over the mountain region in Taiwan, using the WRF model with the Morrison and SBM cloud microphysics schemes. The dust-related immersion freezing parameterization is implemented in the bulk and bin schemes to connect the model with the dust number concentrations in clean and dusty conditions. Multiple datasets including TCCIP precipitation observation, satellite retrieved measurements, and reanalysis are used as references to compare with the model simulations.

The precipitations simulated by the default and modified schemes are shown to be generally consistent with the observations but underestimate the accumulated precipitation and the peak rain rates. The sensitivity simulation results indicate that the long-range transported dust particles can notably affect the convective cloud and precipitation by acting as ice-nucleating particles, even with relatively low number concentrations ($\sim 0.6 \text{ cm}^{-3}$). We find that from clean to dusty conditions, in the heavy precipitation area (High-Pcp), the dust effect enhances the accumulated precipitation by ~15% and ~20% in the precipitation centers and the peak precipitation rate by ~ 41% and ~55% with Morr2 and SBM, respectively.

Both [the](#) Morr2 and SBM schemes predict improved precipitation patterns (pattern correlations increase from 0.56 to 0.6
385 with Morr2 and ~0.55 to 0.66 with SBM) and time series (reduce false predicted rainfall) when dust effect is involved. By
comparing the precipitations simulated by the Morr2 and SBM runs, the result also suggests that the SBM microphysics
scheme predicts [the-a](#) better rainfall intensity, pattern, and time series than the Morr2 scheme in this case.

During the strong convective period (P2, LT 11:00–15:00), the simulated cloud properties (ice/liquid water contents,
cloud area, and cloud top height) are notably impacted due to the dust effects through both convective invigoration and the
390 microphysics effect. ~~The-d~~ Dust aerosols dramatically impact the cloud heterogeneous freezing process during the strong
convective periods, with the ice water content increasing by ~35% and 60% with Morr2 and SBM schemes from clean to
dusty conditions. The increased ice and snow particles in the mixed-phase regime cause the stronger riming and WBF
processes and enhanced latent heating, which invigorate convective intensity and enhance the condensation and melting,
leading to increasing liquid water content and raindrop stronger surface precipitations. During the deep convective period
395 (P2), the mean updraft velocities and latent heating rates averaged over the strong convective grids (over the top 10
percentiles of the updrafts, with $w > 0.5 \text{ m s}^{-1}$) show that dust aerosol effects remarkably increase the condensation and
deposition latent heating rates, leading to the intensified updraft. This study demonstrates the bidirectional influences
between dust aerosols and the convective system, i.e., the low-altitude dust aerosols are lifted to the freezing layer by the
strong convection, and in return, leading the enhanced convective intensity via microphysical and invigoration effects.

400 The result also shows that the Morr2 and SBM schemes simulate different invigoration effects of the dust aerosol, with
a much stronger effect with SBM. From Morr2-Clean to Morr2-Dusty runs, the convective intensity significantly increases at
4–8 km altitude (by ~34%) due to the enhanced condensation latent heating. Meanwhile, with the SBM microphysics
scheme, dust effects intensify the updraft more significantly at upper levels, corresponding to the increase in high-level latent
heating from the deposition process. The different convective invigoration effects of dust simulated by the Morr2 and SBM
405 schemes are indicated partially caused by the saturation adjustment approach used by the bulk scheme. Compared to the
explicit supersaturation forced condensation, the model run with [the](#) saturation adjustment approach predicts [the-a](#) stronger
enhancement of the condensation process at mid-low levels by turning all supersaturated water vapor into liquid water at the
end of each timestep, which limits the deposition process and convective invigoration at upper levels.

In the dust-cloud interaction process, it is difficult to untangle the microphysical effect and convective invigoration
410 effect. With the dust effects, the enhanced freezing and riming can increase latent heat release and boost the convection, and
an increase in latent heating can intensify the updraft and increase condensation and deposition processes in turn. In this
study, the model simulations (SBM, Morr2, and Morr2-ec) averaged over the rain band area demonstrate that both bin and
bulk microphysics schemes can predict the microphysical effects of dust, manifesting as the enhanced immersion freezing
and IWC in the mixed-phase layer. But, among the strong convective grids, the different simulations suggest that the dust
415 effect on the intensity and vertical profile of the convection is sensitive to the supersaturation and condensation approaches
(Morr2 vs. SBM, and Morr2 vs. Morr2-ec) in the microphysics schemes. Besides the saturation adjustment, the different dust

effects predicted by the Morr2 and SBM schemes can also be caused by the different assumptions and parameterizations of the various microphysical processes between the bulk and bin microphysics schemes.

420 This study shows that the dust aerosols transported at low altitudes (below the freezing level) with relatively low
425 number concentrations have remarkable impacts on the orographic convective clouds and precipitation. The extremely low
background dust loading in summer over this region is one possible factor responsible for the significant dust effect, that N_d
increases by about two orders of magnitudes from clean to dusty conditions. The abundant moisture transported by the
circulation could be another influencing factor, which is suggested to be important in the dust-cloud-precipitation
interactions by our previous study (Zhang et al., 2020). This study demonstrates that the importance of dust effects in the
425 predictions of convective cloud and heavy precipitation, and the dust effect improves the simulation by the bulk scheme
which requires less computational resources and is widely used in operational weather forecasting. This study could provide
a possible pointcut to improve the numerical weather prediction of extreme precipitation events.

430 **Data availability.** The GEOS-Chem-APM dust number concentration data can be accessed by contacting the corresponding
authors Yanda Zhang (yzhang31@albany.edu) and Fnagqun Yu (fyu@albany.edu). The MERRA-2 inst3_3d_aer_Nv and
GPM IMERG half-hourly precipitation data are available at Goddard Earth Sciences Data and Information Services Center
(GES DISC) (https://disc.gsfc.nasa.gov/datasets/M2I3NVAER_5.12.4/summary;
https://disc.gsfc.nasa.gov/datasets/GPM_3IMERGHHE_06/summary). The TCCIP precipitation data can be applied and
435 acquired from the Taiwan Climate Change Projection Information and Adaptation Knowledge Platform
(<http://tccip.ncdr.nat.gov.tw/NCDR/main/index.aspx>). MODIS MYD06_L2 data used in this work were acquired from
Atmosphere Archive and Distribution System (LAADS) Distributed Active Archive Center (DAAC)
(https://ladsweb.modaps.eosdis.nasa.gov/search/order/1/MYD06_L2--61)

440 **Author contributions.** YZ, FY, and GL developed the project idea. JF contributed the model code. YZ updated the model.
YZ carried out the numerical simulations. YZ and SL analyzed the data, with contributions from JF. YZ wrote the paper with
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References

- 455 Ault, A. P., Williams, C. R., White, A. B., Neiman, P. J., Creamean, J. M., Gaston, C. J., ... & Prather, K. A.: Detection of Asian dust in California orographic precipitation, *J. Geophys. Res. Atmos.*, 116(D16), doi.org/10.1029/2010JD015351, 2011.
- Andreae, M. O., Charlson, R. J., Bruynseels, F., Storms, H., Van Grieken, R., & Maenhaut, W.: Internal mixture of sea salt, silicates, and excess sulfate in marine aerosols, *Science*, 232(4758), 1620-1623, doi:10.1126/science.232.4758.1620, 1986.
- 460 Atkinson, J., Murray, B., Woodhouse, M. et al. The importance of feldspar for ice nucleation by mineral dust in mixed-phase clouds, *Nature* 498, 355–358, doi:10.1038/nature12278, 2013.
- Bangert, M., Nenes, A., Vogel, B., Vogel, H., Barahona, D., Karydis, V. A., ... & Blahak, U.: Saharan dust event impacts on cloud formation and radiation over Western Europe, *Atmos. Chem. Phys.*, 12(9), 4045-4063. 2012
- Bigg, E. K.: The supercooling of water, *Proc. Phys. Soc. B*, 66(8), 688, doi:10.1088/0370-1301/66/8/309, 1953.
- 465 Chen, C. S., & Chen, Y. L.: The rainfall characteristics of Taiwan, *Mon. Weather Rev.*, 131(7), 1323-1341. doi:10.1175/1520-0493(2003)131%3C1323:TRCOT%3E2.0.CO;2, 2003.
- Chen, J. M., Li, T., & Shih, C. F.: Tropical cyclone–and monsoon-induced rainfall variability in Taiwan, *J. Clim.*, 23(15), 4107-4120. doi:10.1175/2010JCLI3355.1, 2010.
- 470 Chen, J. P., Wang, Z., Young, C. Y., Tsai, F., Tsai, I. C., Wang, G. J., ... & Lu, M. J.: Simulations of Asian Yellow Dust Incursion Over Taiwan for the Spring of 2002 and 2003, *Terr. Atmos. Ocean. Sci.*, 15(5), 949-981, doi:10.3319/TAO.2004.15.5.949(ADSE), 2004.
- Chen, J. P., Hazra, A., Shiu, C. J., Tsai, I. C., & Lee, H. H.: Interaction between aerosols and clouds: current understanding. *Recent Progress in Atmospheric Sciences: Applications to the Asia-Pacific Region*, 231-281, doi:10.1142/9789812818911_0013, 2008.
- 475 Chen, J. P., Yang, C. E., & Tsai, I. C.: Estimation of foreign versus domestic contributions to Taiwan's air pollution, *Atmos. Environ.*, 112, 9-19, doi:10.1016/j.atmosenv.2015.04.022, 2015.

- Chou, C. K., Hsu, W. C., Chang, S. Y., Chen, W. N., Chen, M. J., Huang, W. R., ... & Liu, S. C.: Seasonality of the mass concentration and chemical composition of aerosols around an urbanized basin in East Asia, *J. Geophys. Res. Atmos.*, 122(3), 2026-2042. doi:10.1002/2016JD025728, 2017.
- 480 Chiu, H. F., Tiao, M. M., Ho, S. C., Kuo, H. W., Wu, T. N., & Yang, C. Y.. Effects of Asian dust storm events on hospital admissions for chronic obstructive pulmonary disease in Taipei, Taiwan, *Inhal. Toxicol.*, 20(9), 777-781. doi:10.1080/08958370802005308, 2008.
- Choi, Y. S., Lindzen, R. S., Ho, C. H., & Kim, J.. Space observations of cold-cloud phase change, *Proc. Natl. Acad. Sci. U.S.A.*, 107(25), 11211-11216, doi:10.1073/pnas.1006241107, (2010)
- 485 Guo, J., & Yin, Y. (2015). Mineral dust impacts on regional precipitation and summer circulation in East Asia using a regional coupled climate system model, *J. Geophys. Res. Atmos.*, 120(19), 10-378.
- Carslaw, K. S., Boucher, O., Spracklen, D. V., Mann, G. W., Rae, J. G. L., Woodward, S., and Kulmala, M.: A review of natural aerosol interactions and feedbacks within the Earth system, *Atmos. Chem. Phys.*, 10, 1701–1737, doi:10.5194/acp-10-1701-2010, 2010.
- 490 Choobari, O. A., Zawar-Reza, P., & Sturman, A. (2014). The global distribution of mineral dust and its impacts on the climate system: A review, *Atmos. Res.*, 138, 152-165.
- Creamean, J. M., Suski, K. J., Rosenfeld, D., Cazorla, A., DeMott, P. J., Sullivan, R. C., et al.: Dust and biological aerosols from the Sahara and Asia influence precipitation in the western US. *Science*, 339(6127), 1572–1578. doi:10.1126/science.1227279, 2013.
- 495 Creamean, J. M., Ault, A. P., White, A. B., Neiman, P. J., Ralph, F. M., Minnis, P., and Prather, K. A.: Impact of interannual variations in sources of insoluble aerosol species on orographic precipitation over California's central Sierra Nevada, *Atmos. Chem. Phys.*, 15, 6535–6548, doi:10.5194/acp-15-6535-2015, 2015.
- Creamean, J. M., White, A. B., Minnis, P., Palikonda, R., Spangenberg, D. A., & Prather, K. A.: The relationships between insoluble precipitation residues, clouds, and precipitation over California's southern Sierra Nevada during winter
500 storms, *Atmos. Environ.*, 140, 298-310, doi:10.1016/j.atmosenv.2016.06.016, 2016.
- DeMott, P. J., Sassen, K., Poellot, M. R., Baumgardner, D., Rogers, D. C., Brooks, S. D., et al. (2003). African dust aerosols as atmospheric ice nuclei, *Geophys. Res. Lett.*, 30(14), doi:10.1029/2003gl017410
- DeMott, P. J., Prenni, A. J., Liu, X., Kreidenweis, S. M., Petters, M. D., Twohy, C. H., ... & Rogers, D. (2010). Predicting global atmospheric ice nuclei distributions and their impacts on climate, *Proc. Natl. Acad. Sci. U.S.A.*, 107(25),
505 11217-11222.

- DeMott, P. J., Prenni, A. J., McMeeking, G. R., Sullivan, R. C., Petters, M. D., Tobo, Y., et al.: Integrating laboratory and field data to quantify the immersion freezing ice nucleation activity of mineral dust particles. *Atmos. Chem. Phys.*, 15(1), 393-409. doi:10.5194/acp-15-393-2015, 2015
- 510 [Denjean, C., Cassola, F., Mazzino, A., Triquet, S., Chevaillier, S., Grand, N., ... & Mallet, M., and Formenti, P.: Size distribution and optical properties of mineral dust aerosols transported in the western Mediterranean, *Atmos. Chem. Phys.*, 16, 1081–1104, <https://doi.org/10.5194/acp-16-1081-2016>, 2016.](#)
- Engelstaedter, S., Tegen, I., & Washington, R.: North African dust emissions and transport, *Earth-Sci. Rev.*, 79(1-2), 73–100. doi:10.1016/j.earscirev.2006.06.004, 2006
- European Centre for Medium-Range Weather Forecasts. 2017, updated monthly. ERA5 Reanalysis. Research Data Archive at the National Center for Atmospheric Research, Computational and Information Systems Laboratory, doi:10.5065/D6X34W69. Accessed† 26 10 2020.
- 515
- Fan, J., Leung, L. R., Li, Z., Morrison, H., Chen, H., Zhou, Y., ... & Wang, Y.: Aerosol impacts on clouds and precipitation in eastern China: Results from bin and bulk microphysics, *J. Geophys. Res. Atmos.*, 117(D16). doi:10.1029/2011JD016537, 2012
- 520 Fan, J., Leung, L. R., Rosenfeld, D., Chen, Q., Li, Z., Zhang, J., & Yan, H. Microphysical effects determine macrophysical response for aerosol impacts on deep convective clouds, *Proc. Natl. Acad. Sci. U.S.A.*, 110(48), E4581-E4590, doi:10.1073/pnas.1316830110. 2013.
- Fan, J., Leung, L. R., DeMott, P. J., Comstock, J. M., Singh, B., Rosenfeld, D., et al.: Aerosol impacts on California winter clouds and precipitation during CalWater 2011: local pollution versus long-range transported dust. *Atmos. Chem. Phys.*, 14(1), 81–101. doi:10.5194/acp-14-81-2014, 2014.
- 525
- Fan, J., Leung, L. R., Rosenfeld, D., and DeMott, P. J.: Effects of cloud condensation nuclei and ice nucleating particles on precipitation processes and supercooled liquid in mixed-phase orographic clouds, *Atmos. Chem. Phys.*, 17, 1017–1035, doi:10.5194/acp-17-1017-2017, 2017.
- 530 [Fu, Q., Thorsen, T. J., Su, J., Ge, J. M., & Huang, J. P.: Test of Mie-based single-scattering properties of non-spherical dust aerosols in radiative flux calculations, *J. Quant. Spectrosc. Radiat. Transf.*, 110\(14-16\), 1640-1653, \[doi:10.1016/j.jqsrt.2009.03.010\]\(https://doi.org/10.1016/j.jqsrt.2009.03.010\), 2009.](#)
- Futyan, J. M., & Del Genio, A. D. (2007). Deep convective system evolution over Africa and the tropical Atlantic, *J. Clim.*, 20(20), 5041-5060, doi:10.1175/JAS-D-16-0255.1
- 535 Gatlin, P. N., Petersen, W. A., Knupp, K. R., & Carey, L. D.: Observed response of the raindrop size distribution to changes in the melting layer. *Atmosphere*, 9(8), 319, doi:10.3390/atmos9080319, 2018.

- Gelaro, R., McCarty, W., Suárez, M. J., Todling, R., Molod, A., Takacs, L., ... & Wargan, K.: The modern-era retrospective analysis for research and applications, version 2 (MERRA-2), *J. Clim.*, 30(14), 5419-5454, doi:10.1175/JCLI-D-16-0758.1, 2017.
- 540 Generoso, S., Bey, I., Labonne, M., & Bréon, F. M.: Aerosol vertical distribution in dust outflow over the Atlantic: Comparisons between GEOS-Chem and Cloud-aerosol Lidar and Infrared Pathfinder Satellite Observation (CALIPSO), *J. Geophys. Res. Atmos.*, 113(D24), doi:10.1029/2008JD010154, 2008.
- Gibbons, M., Min, Q., and Fan, J.: Investigating the impacts of Saharan dust on tropical deep convection using spectral bin microphysics, *Atmos. Chem. Phys.*, 18, 12161–12184, doi:10.5194/acp-18-12161-2018, 2018.
- 545 Global Modeling and Assimilation Office (GMAO): MERRA-2 inst3_3d_aer_Nv: 3d,3-Hourly, Instantaneous, Model-Level, Assimilation, Aerosol Mixing Ratio V5.12.4, Greenbelt, MD, USA, Goddard Earth Sciences Data and Information Services Center (GES DISC), Accessed: [2019/09/20], doi: 10.5067/LTVB4GPCOTK2, 2015.
- Grabowski, W. W., & Morrison, H.: Modeling condensation in deep convection, *J. Atmos. Sci.*, 74(7), 2247-2267, doi: 10.1175/JAS-D-16-0255.1, 2017
- 550 Gunn, R., & Phillips, B. B.: An experimental investigation of the effect of air pollution on the initiation of rain. *J. Meteorol.*, 14(3), 272-280, doi:10.1175/1520-0469(1957)014<0272:AEIOTE>2.0.CO;2, 1957.
- Hansen, J., Sato, M., Ruedy, R.: Radiative forcing and climate response, *J. Geophys. Res. Atmos.*, 102 (D6), 6831–6864, doi:10.1029/96JD03436, 1997
- Han, Y., Fang, X., Zhao, T., Bai, H., Kang, S., & Song, L.: Suppression of precipitation by dust particles originated in the Tibetan Plateau, *Atmos. Environ.*, 43(3), 568-574, doi:10.1016/j.atmosenv.2008.10.018, 2009.
- 555 Henny, L., Thorncroft, C. D., Hsu, H. H., & Bosart, L. F.: Extreme Rainfall in Taiwan: Seasonal Statistics and Trends, *J. Clim.*, 34(12), 4711-4731, doi:10.1175/JCLI-D-20-0999.1, 2021
- Hsu, S. C., Huh, C. A., Lin, C. Y., Chen, W. N., Mahowald, N. M., Liu, S. C., ... & Chen, J. P.: Dust transport from non-East Asian sources to the North Pacific, *Geophys. Res. Lett.*, 39(12), doi:10.1029/2012GL051962, 2012.
- 560 [Hoffmann, C., Funk, R., Sommer, M., & Li, Y.: Temporal variations in PM10 and particle size distribution during Asian dust storms in Inner Mongolia. *Atmos. Environ.*, 42\(36\), 8422-8431. doi:10.1016/j.atmosenv.2008.08.014, 2008.](#)
- Hoose, C., & Möhler, O.: Heterogeneous ice nucleation on atmospheric aerosols: a review of results from laboratory experiments, *Atmos. Chem. Phys.*, 12, 9817–9854. doi:10.5194/acp-12-9817-2012, 2012
- Huang, J., Wang, T., Wang, W., Li, Z., & Yan, H.: Climate effects of dust aerosols over East Asian arid and semiarid regions, *J. Geophys. Res. Atmos.*, 119(19), 11-398. doi:10.1002/2014JD021796, 2014.

- 565 Huang, L., Jiang, J. H., Wang, Z., Su, H., Deng, M., & Massie, S. Climatology of cloud water content associated with different cloud types observed by A-Train satellites, *J. Geophys. Res. Atmos.*, 120(9), 4196-4212, doi:10.1002/2014JD022779, 2015.
- Huang, J., Lin, B., Minnis, P., Wang, T., Wang, X., Hu, Y., ... & Ayers, J. K.: Satellite-based assessment of possible dust aerosols semi-direct effect on cloud water path over East Asia, *Geophys. Res. Lett.*, 33(19),
570 doi:10.1029/2006GL026561, 2006
- Huffman, G.J., E.F. Stocker, D.T. Bolvin, E.J. Nelkin, Jackson Tan: GPM IMERG Early Precipitation L3 Half Hourly 0.1 degree x 0.1 degree V06, Greenbelt, MD, Goddard Earth Sciences Data and Information Services Center (GES DISC), Accessed: [2019/06/23], doi:10.5067/GPM/IMERG/3B-HH-E/06, 2019
- Husar, R. B., Tratt, D. M., Schichtel, B. A., Falke, S. R., Li, F., Jaffe, D., et al.: Asian dust events of April 1998, *J. Geophys. Res. Atmos.*, 106(D16), 18,317–18,330. doi:10.1029/2000JD900788, 2001
575
- Iacono, M. J., Delamere, J. S., Mlawer, E. J., Shephard, M. W., Clough, S. A., & Collins, W. D.: Radiative forcing by long-lived greenhouse gases: Calculations with the AER radiative transfer models, *J. Geophys. Res. Atmos.*, doi:10.1029/2008JD009944113(D13). 2008
- Jickells, T. D., An, Z. S., Andersen, K. K., Baker, A. R., Bergametti, G., Brooks, N., ... & Kawahata, H.: Global iron connections between desert dust, ocean biogeochemistry, and climate, *science*, 308(5718), 67-71,
580 doi:10.1126/science.1105959, 2005.
- [Kaaden, N., Massling, A., Schladitz, A., Müller, T., Kandler, K., Schütz, L., ... & Wiedensohler, A.: State of mixing, shape factor, number size distribution, and hygroscopic growth of the Saharan anthropogenic and mineral dust aerosol at Tinfou, Morocco. *Tellus B: Chem. Phys. Meteorol.*, 61\(1\), 51-63. doi: 10.1111/j.1600-0889.2008.00388.x, 2009.](#)
- 585 Khain, A., Lynn, B., & Shpund, J.: High resolution WRF simulations of Hurricane Irene: Sensitivity to aerosols and choice of microphysical schemes, *Atmos. Res.*, 167, 129-145. doi:10.1016/j.atmosres.2015.07.014, 2016
- Khain, A., Pokrovsky, A., Pinsky, M., Seifert, A., and Phillips, V.: Simulation of effects of atmospheric aerosols on deep turbulent convective clouds using a spectral microphysics mixed-phase cumulus cloud model. Part I: Model description and possible applications, *J Atmos Sci*, 61, 2963-2982, doi: 10.1175/Jas-3350.1, 2004.
- 590 Kim, P. S., Jacob, D. J., Fisher, J. A., et al.: Sources, seasonality, and trends of southeast US aerosol: an integrated analysis of surface, aircraft, and satellite observations with the GEOS-Chem chemical transport model, *Atmos. Chem. Phys.*, 15, 10411–10433, doi:10.5194/acp-15-10411-2015, 2015. Klein, H., Nickovic, S., Haunold, W., Bundke, U., Nillius, B., Ebert, M., Weinbruch, S., Schuetz, L., Levin, Z., Barrie, L. A., and Bingemer, H.: Saharan dust and ice nuclei over Central Europe, *Atmos. Chem. Phys.*, 10, 10211–10221, doi:10.5194/acp-10-10211-2010, 2010.

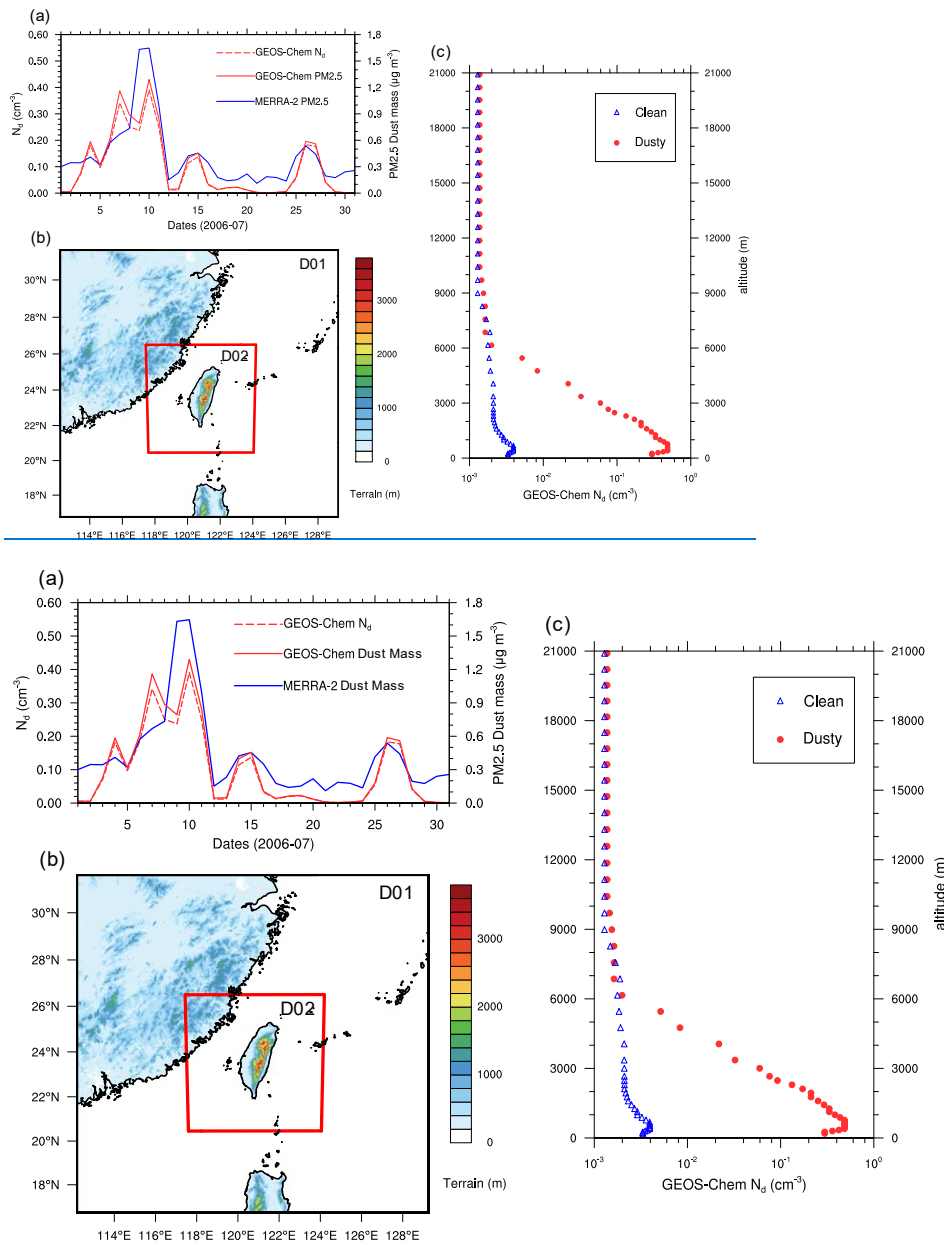
- 595 Knopf, D. A., & Koop, T.: Heterogeneous nucleation of ice on surrogates of mineral dust, *J. Geophys. Res. Atmos.*, 111(D12), doi:10.1029/2005JD006894, 2006.
- Kok, J., Ridley, D., Zhou, Q. et al.: Smaller desert dust cooling effect estimated from analysis of dust size and abundance, *Nat. Geosci.*, 10, 274–278, doi:10.1038/ngeo2912, 2017.
- 600 Koren, I., Kaufman, Y.J., Remer, L.A., Kaufman, Y.J., Remer, L.A., Koren, I., Martins, J.V., Martins, J.V.: Measurement of the effect of Amazon smoke on inhibition of cloud formation. *Science* 303 (5662), 1342–1345, doi:10.1126/science.1089424, 2004.
- Lebo, Z. J., Morrison, H., and Seinfeld, J. H.: Are simulated aerosol-induced effects on deep convective clouds strongly dependent on saturation adjustment?, *Atmos. Chem. Phys.*, 12, 9941–9964, doi:10.5194/acp-12-9941-2012, 2012.
- 605 Lebo, Z. J. and Seinfeld, J. H.: Theoretical basis for convective invigoration due to increased aerosol concentration, *Atmos. Chem. Phys.*, 11, 5407–5429, doi:10.5194/acp-11-5407-2011, 2011.
- Levin, Z., Teller, A., Ganor, E., & Yin, Y.: On the interactions of mineral dust, sea-salt particles, and clouds: A measurement and modeling study from the Mediterranean Israeli Dust Experiment campaign, *J. Geophys. Res. Atmos.*, 110(D20), doi.org/10.1029/2005JD005810, 2005.
- Li, X., Maring, H., Savoie, D., Voss, K., & Prospero, J. M.: Dominance of mineral dust in aerosol light-scattering in the
610 North Atlantic trade winds, *Nature*, 380(6573), 416–419. doi:10.1038/380416a0, 1996.
- [Li Z., Y. Wang, J. Guo, et al.: East Asian study of tropospheric aerosols and their impact on regional clouds, precipitation, and climate \(EASTAIR\(CPC\)\). *Journal of Geophysical Research: Atmospheres*. 124 \(23\), 13026-13054. doi: 10.1029/2019JD030758, 2019.](#)
- 615 Lin, C. Y., Chou, C. C., Wang, Z., Lung, S. C., Lee, C. T., Yuan, C. S., ... & Liu, S. C.: Impact of different transport mechanisms of Asian dust and anthropogenic pollutants to Taiwan, *Atmos. Environ.*, 60, 403–418, doi:10.1016/j.atmosenv.2012.06.049, 2012.
- Liu, Z., Omar, A., Vaughan, M., Hair, J., Kittaka, C., Hu, Y., ... & Pierce, R. (2008). CALIPSO lidar observations of the optical properties of Saharan dust: A case study of long-range transport, *J. Geophys. Res. Atmos.*, 113(D7), doi:10.1029/2007JD008878, 2008.
- 620 [Mahowald, N., Albani, S., Kok, J. F., Engelstaeder, S., Scanza, R., Ward, D. S., & Flanner, M. G.: The size distribution of desert dust aerosols and its impact on the Earth system. *Aeolian Res.*, 15, 53-71, doi:10.1016/j.aeolia.2013.09.002, 2014.](#)
- Meyers, M. P., DeMott, P. J., & Cotton, W. R.: New primary ice-nucleation parameterizations in an explicit cloud model, *J. Appl. Meteorol.*, 31(7), 708–721, doi:10.1175/1520-0450(1992)031<0708:npinpi>2.0.co;2, 1992.

- 625 Min, Q.-L., Li, R., Lin, B., Joseph, E., Wang, S., Hu, Y., Morris, V., and Chang, F.: Evidence of mineral dust altering cloud microphysics and precipitation, *Atmos. Chem. Phys.*, 9, 3223–3231, doi:10.5194/acp-9-3223-2009, 2009.
- Morrison, H., Thompson, G., & Tatarskii, V.: Impact of cloud microphysics on the development of trailing stratiform precipitation in a simulated squall line: Comparison of one-and two-moment schemes, *Mon. Weather Rev.*, 137(3), 991-1007, doi:10.1175/2008MWR2556.1, 2009.
- 630 Muhlbauer, A., & Lohmann, U.: Sensitivity studies of aerosol–cloud interactions in mixed-phase orographic precipitation, *J. Atmos. Sci.*, 66(9), 2517-2538. doi:10.1175/2009JAS3001.1, 2009.
- [Penner, J. E., Andreae, M. O., Annegarn, H., Barrie, L., Feichter, J., Hegg, D., ... & Pitari, G.: Aerosols, their direct and indirect effects. In *Climate Change 2001: The Scientific Basis. Contribution of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change* \(pp. 289-348\). Cambridge University Press, 2001.](#)
- 635
- Perry, K. D., Cliff, S. S., & Jimenez-Cruz, M. P.: Evidence for hygroscopic mineral dust particles from the Intercontinental Transport and Chemical Transformation Experiment, *J. Geophys. Res. Atmos.*, 109(D23), doi:10.1029/2004JD004979, 2004.
- Platnick, S., S. A. Ackerman, M. D. King, K. Meyer, W. P. Menzel, R. E. Holz, B. A. Baum, and P. Yang: MODIS atmosphere L2 cloud product (06_L2), NASA MODIS Adaptive Processing System, Goddard Space Flight Center. doi:10.5067/MODIS/MYD06_L2.006, 2015.
- 640
- Pope, C.A., Burnett, R.T., Thun, M.J., Calle, E.E., Krewski, D., Ito, K., Thurston, G.D.: Lung cancer, cardiopulmonary mortality, and long-term exposure to fine particulate air pollution. *J. Am. Med. Assoc.* 287 (9), 1132–1141. doi:10.1001/jama.287.9.1132, 2002.
- 645 Porz, N., Hanke, M., Baumgartner, M., & Spichtinger, P.: A model for warm clouds with implicit droplet activation, avoiding saturation adjustment. *Mathematics of Climate and Weather Forecasting*, 4(1), 50-78. doi:10.1515/mcwf-2018-0003, 2018.
- Prospero, J.M.: Long-term measurements of the transport of African mineral dust to the southeastern United States: Implications for regional air quality. *J. Geophys. Res. Atmos.*, 104 (D13), 15917–15927, 1999.
- 650 Pruppacher, H. R. and Klett, J. D.: *Microphysics of Clouds and Precipitation*, Kluwer Academic Publishers, Boston, 1997.
- Ridley, D. A., Heald, C. L., Kok, J. F., & Zhao, C. (2016). An observationally constrained estimate of global dust aerosol optical depth, *Atmos. Chem. Phys.*, 16(23), 15097-15117. doi:10.5194/acp-16-15097-2016
- Roberts, P., & Hallett, J.: A laboratory study of the ice nucleating properties of some mineral particulates, *Q. J. R. Meteorol. Soc.*, 94(399), 25-34. doi:10.1002/qj.49709439904, 1968

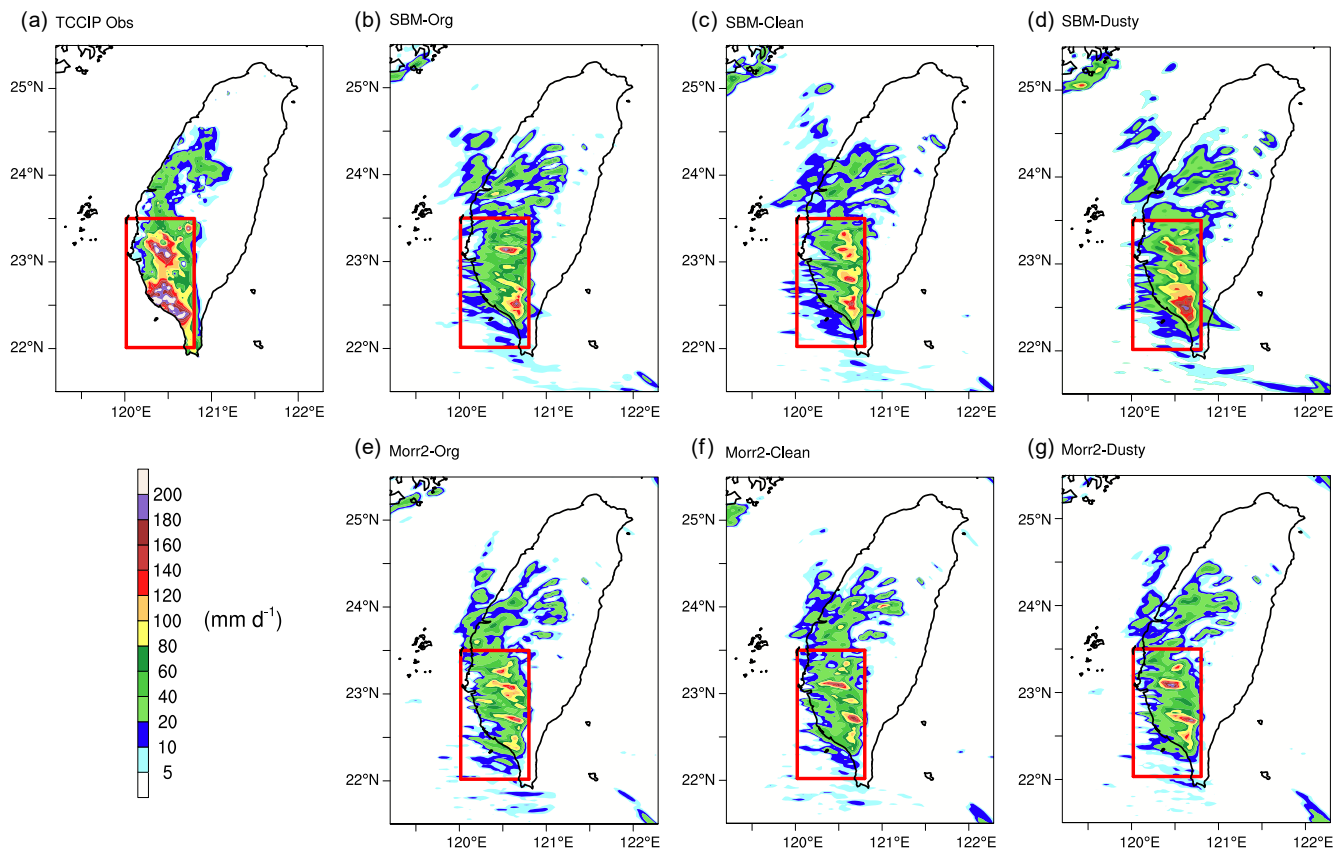
- 655 Rosenfeld, D., Rudich, Y., & Lahav, R.: Desert dust suppressing precipitation: A possible desertification feedback loop, *Proc. Natl. Acad. Sci. U.S.A.*, 98(11), 5975-5980, 2001.
- Sassen, K., DeMott, P. J., Prospero, J. M., & Poellot, M. R.: Saharan dust storms and indirect aerosol effects on clouds: CRYSTAL-FACE results, *Geophys. Res. Lett.*, 30(12), doi:10.1029/2003GL017371, 2003
- Seigel, R. B., Van Den Heever, S. C., & Saleeby, S. M.: Mineral dust indirect effects and cloud radiative feedbacks of a
660 simulated idealized nocturnal squall line, *Atmos. Chem. Phys.*, 13(8), doi:10.5194/acp-13-4467-2013, 2013.
- Si, M., Irish, V. E., Mason, R. H., Vergara-Temprado, J., Hanna, S. J., Ladino, L. A., Yakobi-Hancock, J. D., Schiller, C. L., Wentzell, J. J. B., Abbatt, J. P. D., Carslaw, K. S., Murray, B. J., and Bertram, A. K.: Ice-nucleating ability of aerosol particles and possible sources at three coastal marine sites, *Atmos. Chem. Phys.*, 18, 15669–15685, <https://doi.org/10.5194/acp-18-15669-2018>, 2018.
- 665 Skamarock, W. C., J. B. Klemp, J. Dudhia, D. O. Gill, Z. Liu, J. Berner, W. Wang, J. G. Powers, M. G. Duda, D. M. Barker, and X.-Y. Huang: A Description of the Advanced Research WRF Version 4. NCAR Tech. Note NCAR/TN-556+STR, 145 pp, doi:10.5065/1dfh-6p97, 2019
- Smoydzin, L., Teller, A., Tost, H., Fnais, M., & Lelieveld, J: Impact of mineral dust on cloud formation in a Saharan outflow region, *Atmos. Chem. Phys.*, 12(23), 2012.
- 670 Solomos, S., Kallos, G., Kushta, J., Astitha, M., Tremback, C., Nenes, A., and Levin, Z.: An integrated modeling study on the effects of mineral dust and sea salt particles on clouds and precipitation, *Atmos. Chem. Phys.*, 11, 873–892, doi:10.5194/acp-11-873-2011, 2011.
- Storer, R. L., Van den Heever, S. C., & L'Ecuyer, T. S.: Observations of aerosol-induced convective invigoration in the tropical east Atlantic, *J. Geophys. Res. Atmos.*, 119(7), 3963-3975, 2014.
- 675 Tiedtke, M.: A comprehensive mass flux scheme for cumulus parameterization in large-scale models, *Mon. Weather Rev.*, 117(8), 1779-1800, doi:10.1175/1520-0493(1989)117<1779:ACMFSF>2.0.CO;2, 1989.
- Teller, A. and Levin, Z.: The effects of aerosols on precipitation and dimensions of subtropical clouds: a sensitivity study using a numerical cloud model, *Atmos. Chem. Phys.*, 6, 67–80, doi:10.5194/acp-6-67-2006, 2006.
- Van den Heever, S. C., Carrió, G. G., Cotton, W. R., DeMott, P. J., & Prenni, A. J.: Impacts of nucleating aerosol on Florida
680 storms. Part I: Mesoscale simulations, *J. Atmos. Sci.*, 63(7), 1752-1775, doi:10.1175/JAS3713.1, 2006.
- Textor, C., Schulz, M., Guibert, S., Kinne, S., Balkanski, Y., Bauer, S., ... & Dentener, F.: Seland; Stier, P.; Takemura, T.; Tie, X. Analysis and quantification of the diversities of aerosol life cycles within AeroCom, *Atmos. Chem. Phys.*, 6, 1777-1813, doi:10.5194/acp-6-1777-2006, 2006.

- 685 Uno, I., Eguchi, K., Yumimoto, K., Takemura, T., Shimizu, A., Uematsu, M., ... & Sugimoto, N.: Asian dust transported one full circuit around the globe. *Nature Geoscience*, 2(8), 557. doi:10.1038/ngeo583, 2009
- [Wang, W., Huang, J., Minnis, P., Hu, Y., Li, J., Huang, Z., ... & Wang, T.: Dusty cloud properties and radiative forcing over dust source and downwind regions derived from A-Train data during the Pacific Dust Experiment. *J. Geophys. Res. Atmos.*, 115\(D4\). doi:10.1029/2010JD014109, 2010.](#)
- 690 Wang, Y., Fan, J., Zhang, R., Leung, L. R., & Franklin, C.: Improving bulk microphysics parameterizations in simulations of aerosol effects, *J. Geophys. Res. Atmos.*, 118(11), 5361-5379. doi:10.1002/jgrd.50432, 2013
- Xu, J. W., Martin, R. V., Morrow, A., Sharma, S., Huang, L., Leaitch, W. R., et al.: Source attribution of Arctic black carbon constrained by aircraft and surface measurements. *Atmos. Chem. Phys.*, 17(19), 11971-11989. doi:10.5194/acp-17-11971-2017, 2017.
- Xu, W., Zipser, E. J., & Liu, C.: Rainfall characteristics and convective properties of mei-yu precipitation systems over South China, Taiwan, and the South China Sea. Part I: TRMM observations. *Monthly Weather Review*, 137(12), 4261-4275. doi:10.1175/2009MWR2982.1, 2009.
- 695 Yim, S. Y., Wang, B., Xing, W., & Lu, M. M.: Prediction of Meiyu rainfall in Taiwan by multi-lead physical–empirical models. *Climate Dynamics*, 44(11-12), 3033-3042. doi:10.1007/s00382-014-2340-0, 2015.
- Yin, Y., Wurzler, S., Levin, Z., & Reisin, T. G.: Interactions of mineral dust particles and clouds: Effects on precipitation and cloud optical properties, *J. Geophys. Res. Atmos.*, 107(D23), AAC-19, doi:10.1029/2001JD001544, 2002.
- 700 Zender, C.S., Bian, H., Newman, D.: Mineral Dust Entrainment and Deposition (DEAD) model: Description and 1990s dust climatology, *J. Geophys. Res. Atmos.*, 108 (D14), 4416. doi:10.1029/2002JD002775, 2003.
- Zender, C.S., Miller, R.L., Tegen, I., 2004. Quantifying mineral dust mass budgets: Terminology, constraints, and current estimates, *EOS Trans.* 85 (48), 509–512.
- 705 Zimmermann, F., Weinbruch, S., Schütz, L., Hofmann, H., Ebert, M., Kandler, K., & Worringer, A.: Ice nucleation properties of the most abundant mineral dust phases, *J. Geophys. Res. Atmos.*, 113(D23). doi:10.1029/2008JD010655s, 2008
- Zhang, H., McFarquhar, G. M., Saleeby, S. M., & Cotton, W. R.: Impacts of Saharan dust as CCN on the evolution of an idealized tropical cyclone, *Geophys. Res. Lett.*, 34(14), doi:10.1029/2007GL029876, 2007.
- 710 Zhao, C., Liu, X., Ruby Leung, L., and Hagos, S.: Radiative impact of mineral dust on monsoon precipitation variability over West Africa, *Atmos. Chem. Phys.*, 11, 1879–1893, doi:10.5194/acp-11-1879-2011, 2011.
- Zhang, Y., Luo, G., & Yu, F.: Seasonal Variations and Long-Term Trend of Dust Particle Number Concentration Over the Northeastern United States, *J. Geophys. Res. Atmos.*, 124(23), 13140-13155, doi:10.1029/2019JD031388, 2019.

- 715 Zhang, Y., Cai, Y.J., Yu, F., Luo, G., Chou, C.C.K.: Seasonal Variations and Long-term Trend of Mineral Dust Aerosols
over the Taiwan Region. *Aerosol Air Qual. Res.* doi:10.4209/aaqr.2020.07.0433, 2020a.
- Zhang, Y., Yu, F., Luo, G., Chen, J. P., & Chou, C. C. K.: Impact of mineral dust on summertime precipitation over the
Taiwan region, *J. Geophys. Res. Atmos.*, 125(19), e2020JD033120. doi:10.1029/2020JD033120, 2020b.
- Zhang, Y., Fan, J., Li, Z., and Rosenfeld, D.: Impacts of cloud microphysics parameterizations on simulated aerosol–cloud
interactions for deep convective clouds over Houston, *Atmos. Chem. Phys.*, 21, 2363–2381, doi:10.5194/acp-21-
720 2363-2021, 2021.



725 **Figure 1.** (a) Time series of the GEOS-Chem-APM simulated dust number concentration (N_d : red dashed line) and [PM_{2.5}](#) dust [mass](#) mixing ratio (red solid line) and MERRA-2 [PM_{2.5}](#)-dust [mass](#) mixing ratio reanalysis (blue line) in Jul. 2006, (b) simulation domains and terrain height, and (c) vertical profiles of WRF boundary conditions of dust number concentration in Clean (blue triangles) and Dusty (red dots) runs.



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Figure 2. 24-hour accumulated precipitation (Γ_{24h}) (LT 2006-07-08) observations from (a) TCCIP, and simulations by (b) SBM-Org, (c) SBM-Clean, (d) SBM-Dusty, (e) Morr2-Org, (f) Morr2-Clean, and (g) Morr2-Dusty. The heavy precipitation is mainly located in the area marked by the red box (High-Pcp).

24h Precipitation (LT:2006-07-08)

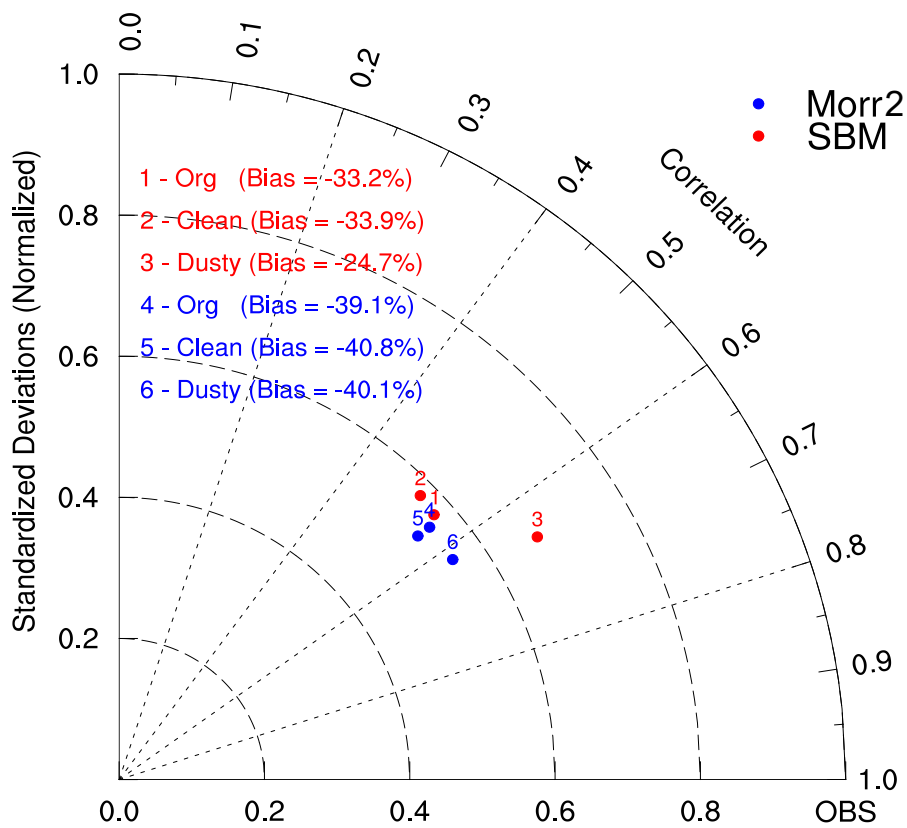
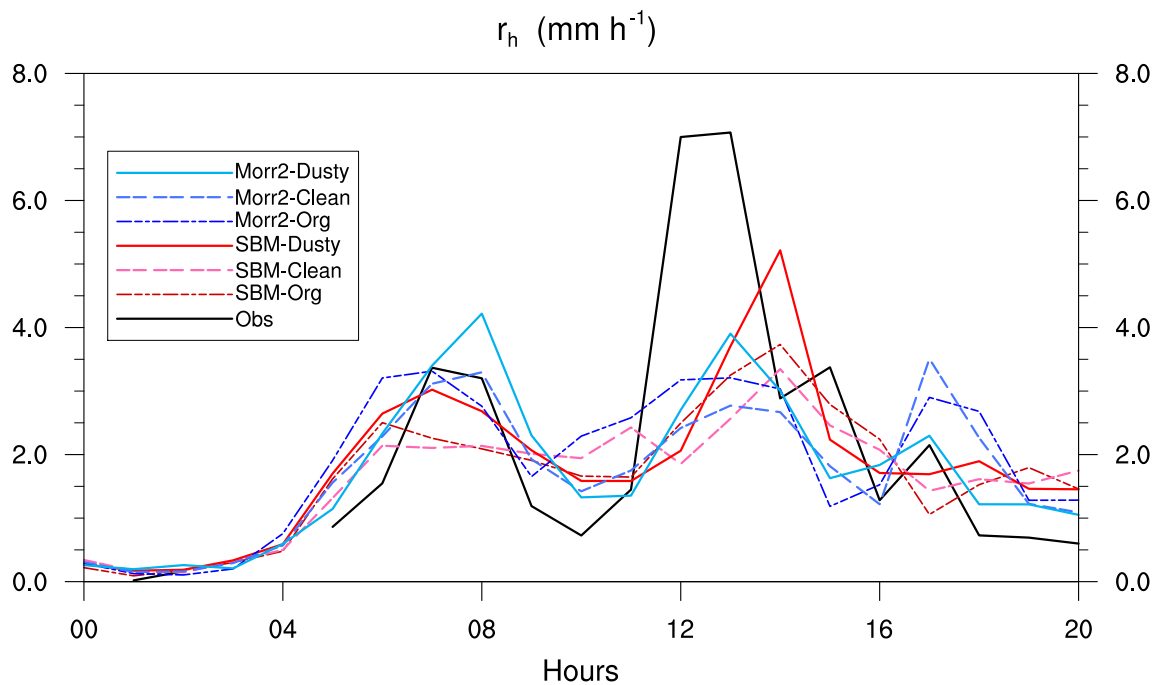
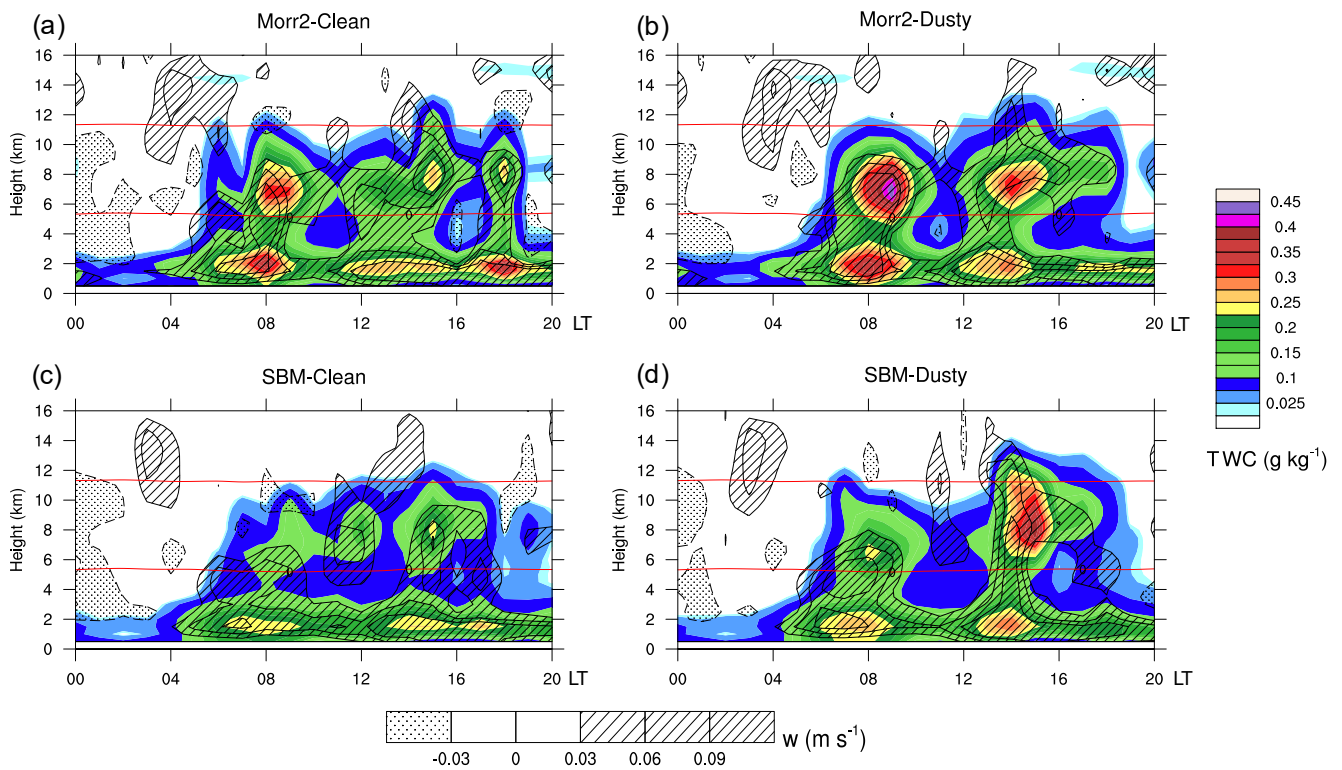


Figure 3. Taylor diagram showing spatial correlations coefficients (r), standard deviations (SD), and bias (%) between simulations by the six numerical runs and the observation of the 24-h accumulated precipitation.

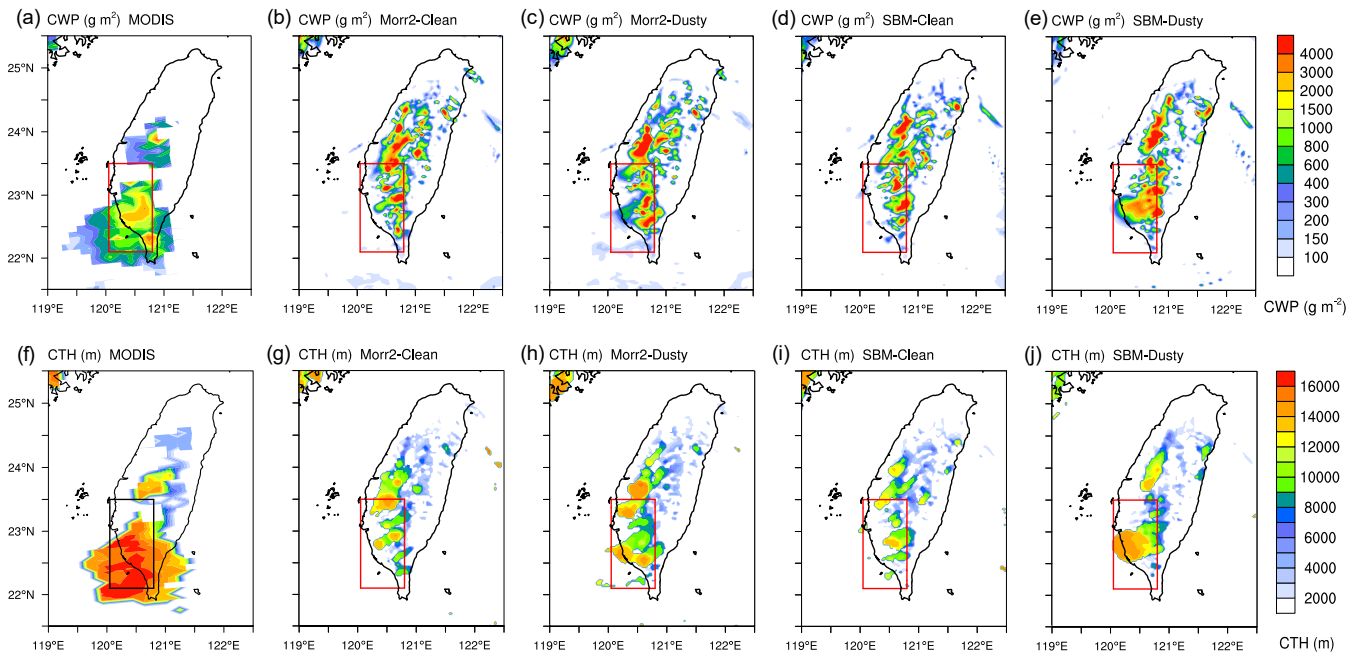


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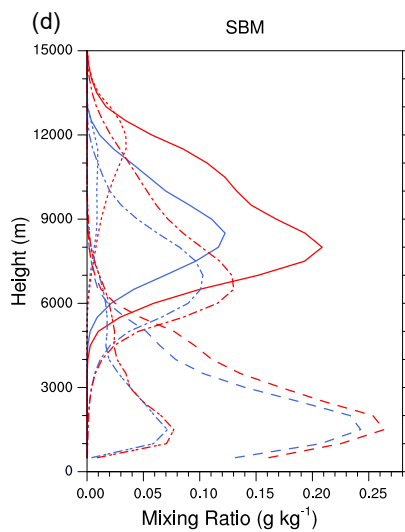
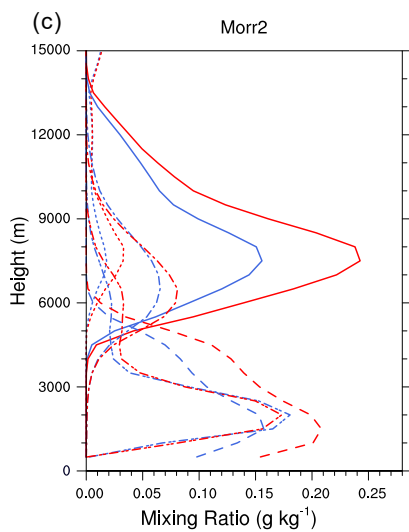
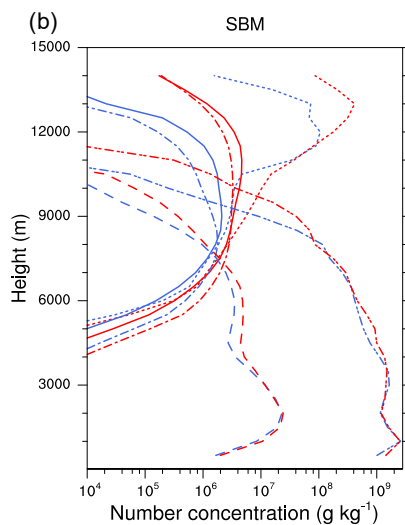
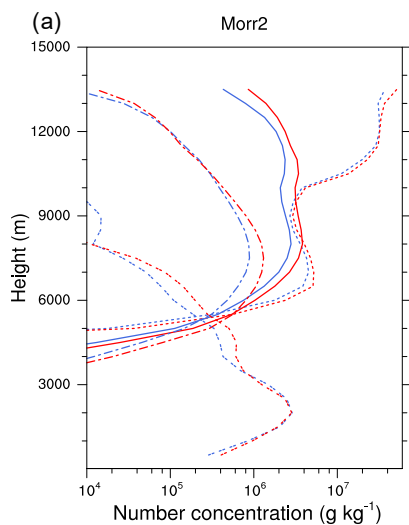
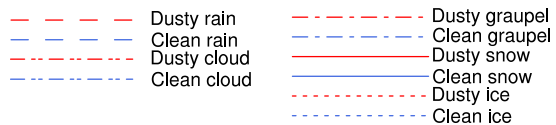
Figure 4. Hourly rainfall rates from IMERG observation (black line) and six numerical runs (colored lines) (LT 00:00–20:00, 08 July 2006), averaged among rainfall grids ($> 0 \text{ mm h}^{-1}$) in the High-Pcp area.



745 **Figure 5.** (a-d) Time series of total water content TWC (shaded) and vertical velocity (contours) averaged among rainfall grids ($> 0 \text{ mm h}^{-1}$) in the High-Pcp area for model simulations of Morr2-Clean, Morr2-Dusty, SBM-Clean, and SBM-Dusty, respectively. The two red lines denote the temperature of $0 \text{ }^{\circ}\text{C}$ and $-38 \text{ }^{\circ}\text{C}$.



750 **Figure 6.** The cloud water path (CWP) at LT 14:00 from (a) MODIS observation, (b) Morr2-Clean, (c) Morr2-Dusty, (d) SBM-Clean, (e) SBM-Dusty simulations. Cloud top height (CTH) on grids with CWP larger than 100 g m^{-2} from (f) MODIS observation, (g) Morr2-Clean, (h) Morr2-Dusty, (i) SBM-Clean, and (j) SBM-Dusty simulations.



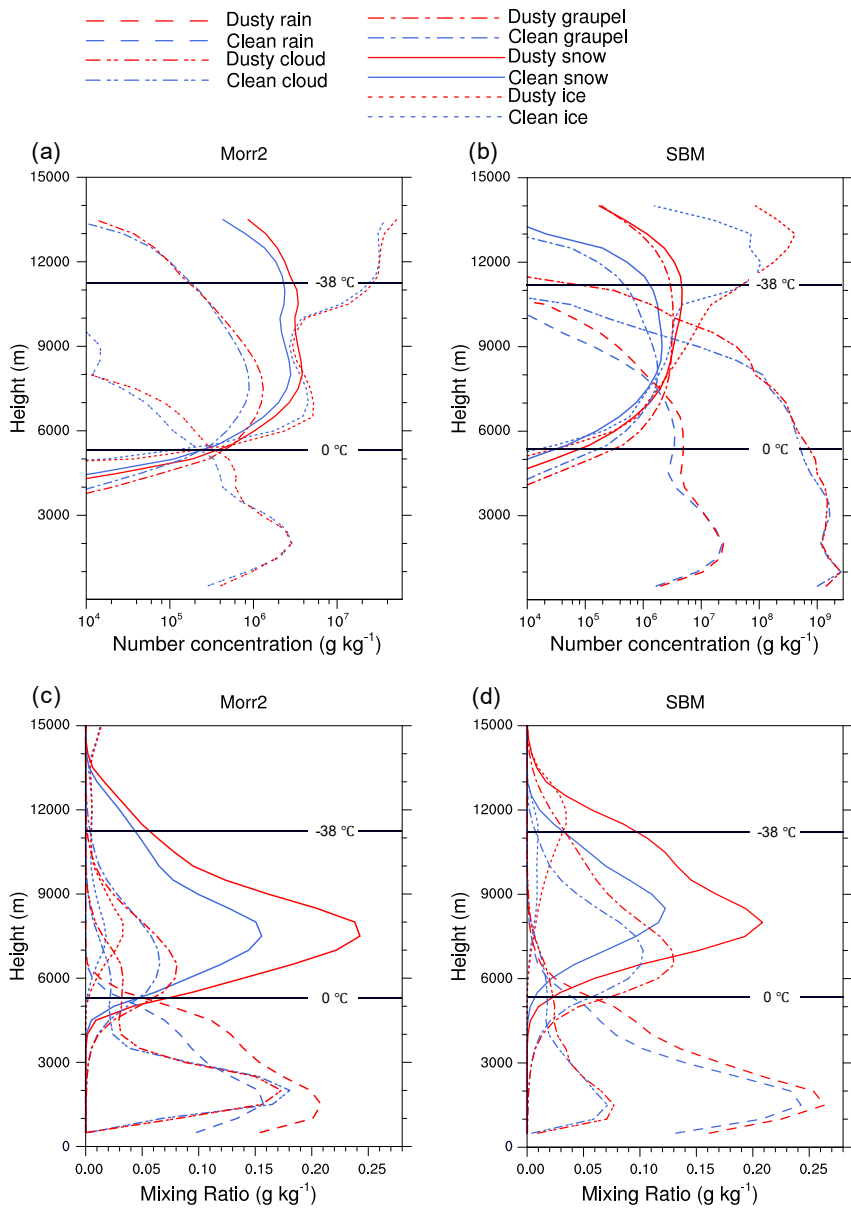


Figure 7. Vertical profiles of number concentrations of cloud hydrometers averaged among precipitation grids in the High-Pcp area from LT 11:00–15:00 simulated by (a) Morr2-Clean (blue lines) & Dusty (red lines), and (b) SBM-Clean (blue lines) & Dusty (red lines) runs. Similar time-space averaged vertical profiles of hydrometeor mass mixing ratios by (c) Morr2-Clean (blue lines) & Dusty (red lines), and (d) SBM-Clean (blue lines) & Dusty (red lines).

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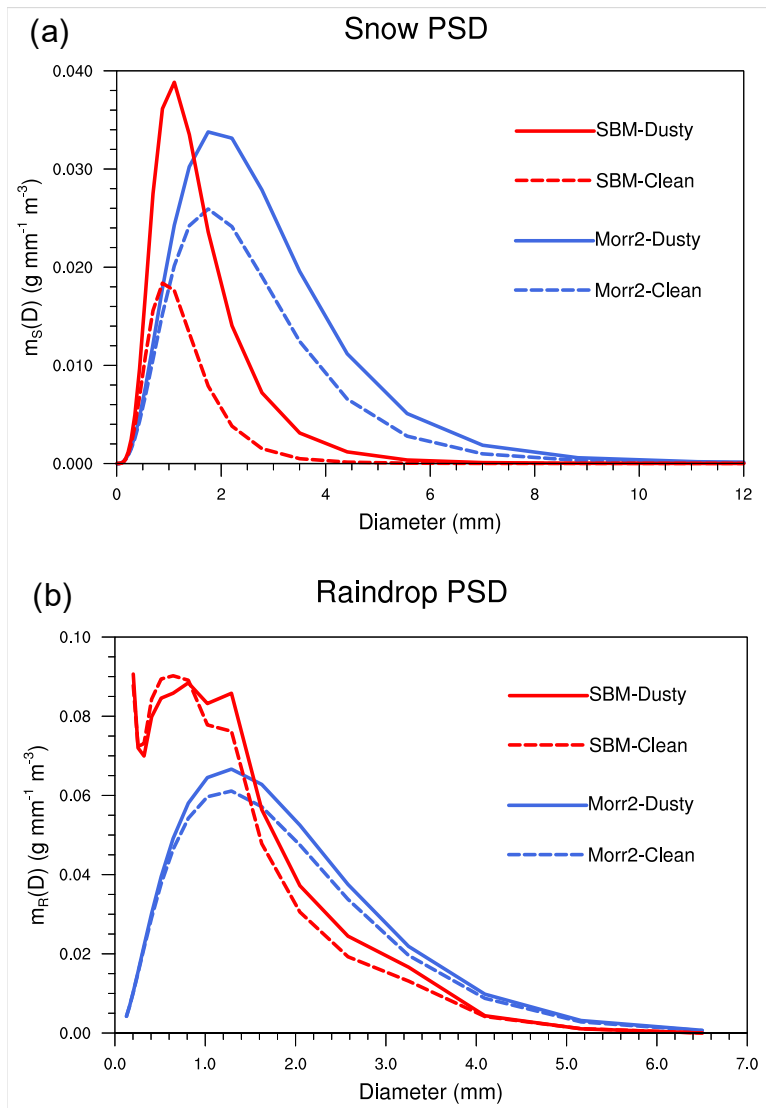


Figure 8. Particle mass spectrum size distributions of the (a) snow particle and (b) raindrop from SBM-Clean (red dashed line) & Dusty (red solid line) and Morr2-Clean (blue dashed line) and Dusty (blue solid line) runs averaged over all the precipitation grids in the High-Pcp area during P2.

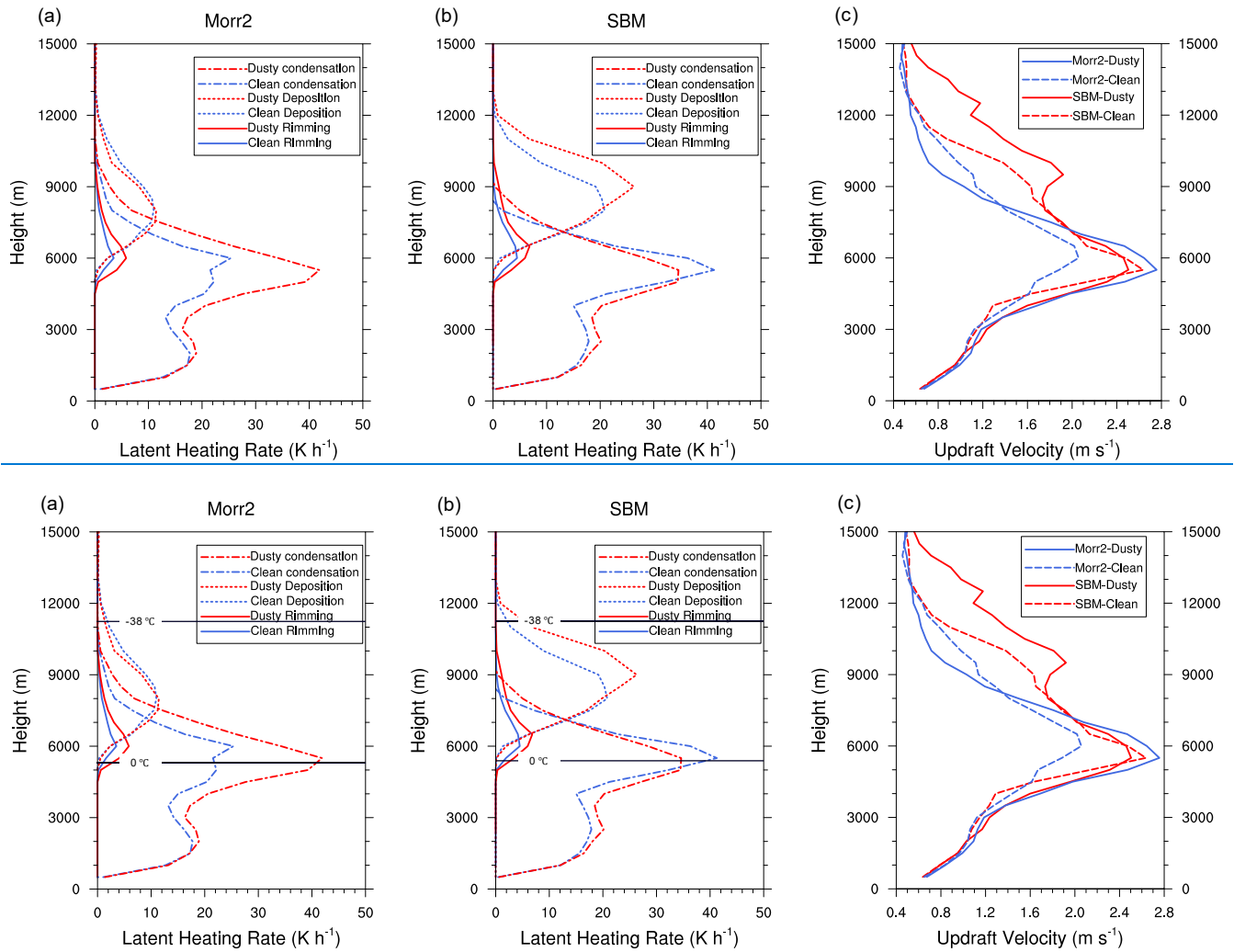
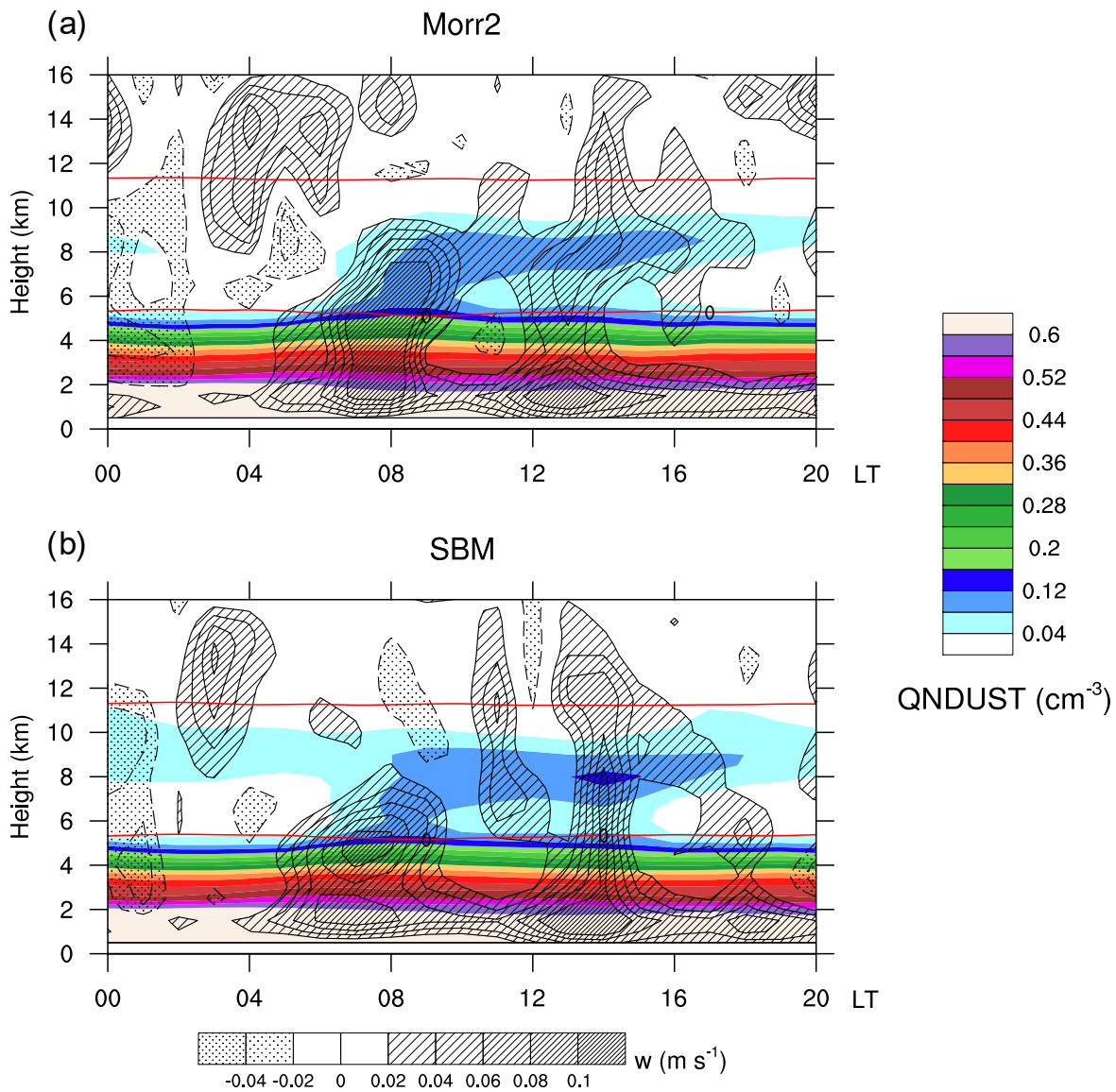
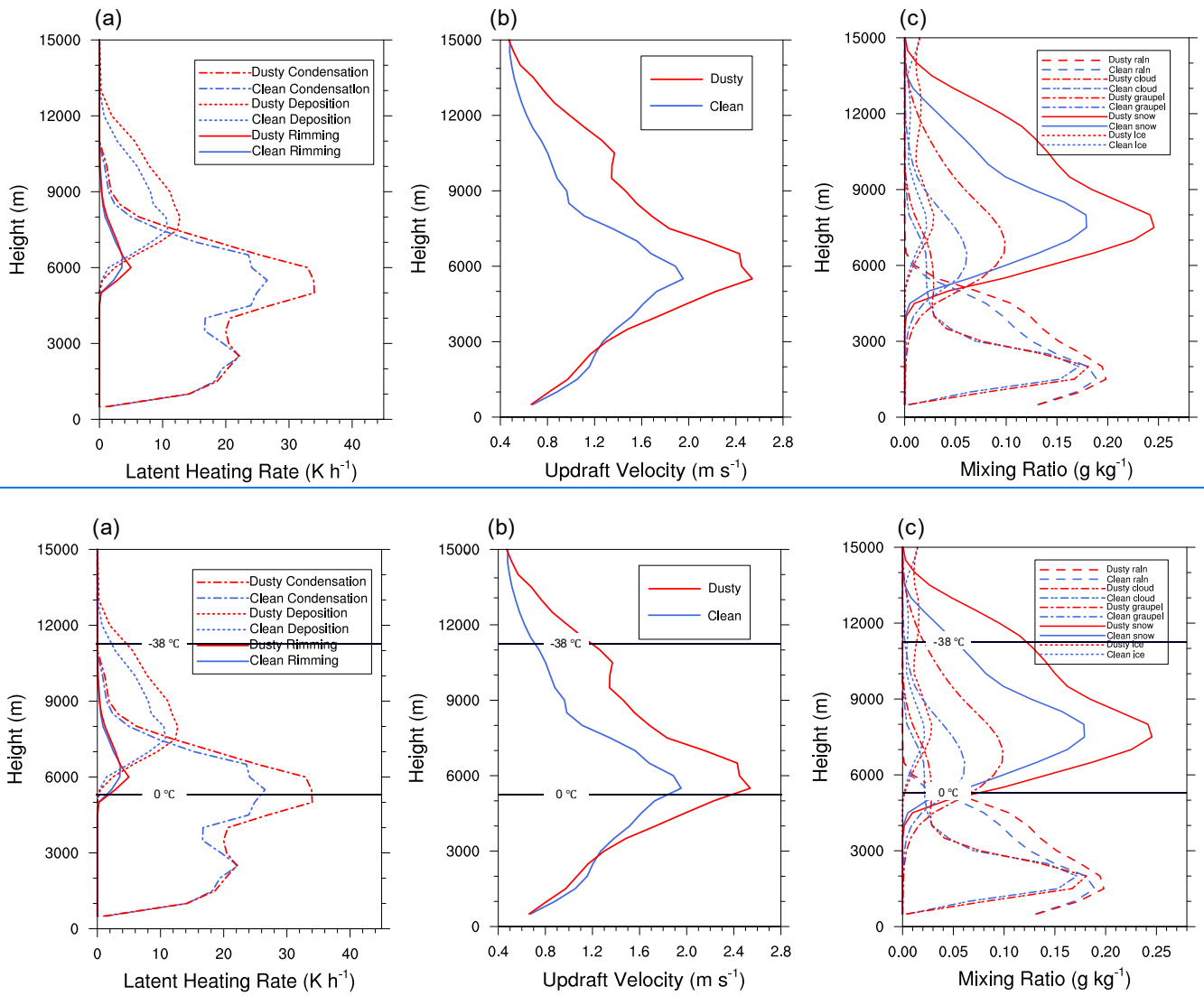


Figure 9. (a) Vertical profiles of latent heating rates averaged over the top 10 percentiles of the updrafts ($w > 0.5 \text{ m s}^{-1}$) from (ab) Morr2-Clean (blue lines) & Dusty (red lines), and (eb) SBM-Clean (blue lines) & Dusty (red lines) in High-Pcp area during LT 12:00-16:00 (P2). (c) Similar time-space averaged vertical profiles of updraft velocity averaged over the top 10 percentiles of the updrafts (90th to 100th, with $w > \sim 0.5 \text{ m s}^{-1}$) simulated by Morr2-Clean & Dusty (blue) and SBM-Clean & Dusty (red).

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775 **Figure 10.** Time series of dust number concentration (shaded) and temperature (red line) and vertical velocity (contours) averaged among precipitation grids (> 0 mm h⁻¹) in the High-Pcp area from (b) Morr2-Dusty and (c) SBM-Dusty.



780 **Figure 11.** Vertical profiles of (a) latent heating rates and (b) vertical velocities averaged over the top 10 percentiles of the updrafts (90th to 100th, with $w > \sim 0.5 \text{ m s}^{-1}$) in High-Pcp area during P2 (LT 12:00–16:00) from Morr2-ec-Clean (blue lines) & Morr2-ec-Dusty (red lines). (c) Vertical profiles of cloud hydrometer mixing ratios averaged among precipitation grids in the High-Pcp area during P2 simulated by Morr2-ec-Clean (blue lines) & Morr2-ec-Dusty (red lines).

785 **Table 1.** Six numerical runs with different settings of the immersion freezing parameterization are considered in this study. The default Morrison and SBM schemes in WRFv4.2.0 with Bigg (1953) immersion freezing parameterization are used as control runs (Morr2-Org and SBM-Org). Four simulations with modified Morr2 and SBM schemes were run with different dust profiles simulated by the GEOS-Chem model, respectively.

Microphysics scheme	Experiment runs	Immersion freezing parameterization
Morr2	Morr2-Org	Bigg (1953)
	Morr2-Clean	DeMott et al. (2015)
	Morr2-Dusty	DeMott et al. (2015)
SBM	SBM-Org	Bigg (1953)
	SBM-Clean	DeMott et al. (2015)
	SBM-Dusty	DeMott et al. (2015)

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Table 2. Cases to investigate the influences of the saturation adjustment on the dust-cloud interaction, based on the settings of Morr2-Clean & Dusty. The saturation adjustment assumption is replaced by the explicit calculation of the water vapor surplus/condensation (Morr2-ec-Clean & Dusty).

<u>Experiment runs</u>	<u>Change of the water vapor surplus</u>
<u>Morr2-Clean</u>	<u>Saturation adjustment</u>
<u>Morr2-Dusty</u>	<u>Saturation adjustment</u>
<u>Morr2-ec-Clean</u>	<u>Explicit calculation based on Eq. (2)</u>
<u>Morr2-ec- Dusty</u>	<u>Explicit calculation based on Eq. (2)</u>