

In this work, the author's present an in-depth look at the gas-particle partitioning of a handful of important, highly oxygenated compounds. Uncertainty in the partitioning of these compounds, and the potential of salting-in to describe observed discrepancies is an important unanswered question that is of interest to readers of this journal. Overall, this work is technically sound and should be published. Some minor concerns and comments are described below:

Response:

Thanks for the reviewer's comments, and we'll reply these point by point.

General comments:

1. The approach to modeling partitioning nicely accounts for partitioning between phases, but a lot of the information to understand their approach is split between the SI and the main text. Some of the SI I think should maybe be brought into the main text (probably at least Eq. S1 and/or S2, and Figure S2)

Response:

Thanks. We have moved those equations and Figure S2 (now Figure 1) into the main text. To clarify the approach used for the exploration of gas-particle partitioning of polyol tracers, three partitioning cases were defined in the revised manuscript.

Pages 10-13, lines 238-295

“Calculations of partitioning coefficients. Here, we defined three partitioning cases to explore the influence of dissolution in aerosol liquid water on gas-particle partitioning of polyol tracers in the atmosphere. *Case 1* presumes instantaneous equilibrium between the gas phase and particulate OM based on the equilibrium absorptive partitioning theory. In this case, particulate OM is assumed to be the only absorbing phase and behave as an ideal solution. Then the absorptive gas-particle partitioning coefficients ($K_{p,OM}$, $m^3 \mu g^{-1}$) were calculated from measurements ($K_{p,OM}^m$) and predicted theoretically ($K_{p,OM}^t$) as follows

$$K_{p,OM}^m = \frac{F/M_{OM}}{A} \quad (2)$$

$$K_{p,OM}^t = \frac{RT}{10^6 \overline{MW}_{OM} \zeta_{OM} p_L^\circ} \quad (3)$$

where M_{OM} denotes the mass concentration of absorptive organic matter (OM = OC \times 1.6; Turpin and Lim, 2001); F ($ng\ m^{-3}$) and A ($ng\ m^{-3}$) are particulate and gaseous concentrations of individual polyols, respectively. In eq 3, R ($m^3\ atm\ K^{-1}\ mol^{-1}$) and T (K) are the ideal gas constant and ambient temperature; \overline{MW}_{OM} , average molecular weight of absorptive OM, is set at $200\ g\ mol^{-1}$ for all samples (Barsanti and Pankow, 2004; Williams et al., 2010); ζ_{OM} denotes the mole fraction scale activity coefficient, and is presumed to be unity for all species in each sample; p_L° (atm) is the vapor pressure of each pure compound, and is predicted with several estimation tools and adjusted for each sampling interval based on the average temperature (Text S3 and Table S4).

Due to the influence of mixing state and water content in aerosols, several studies modeled the gas-particle partitioning of oxygenated organic compounds by defining a liquid-liquid phase separation (LLPS) in the aerosol (Zuend and Seinfeld, 2012; Pye

et al., 2018). The organic-inorganic interactions and changes of activity coefficients in aqueous mixtures were fully considered as well. In this study, we proposed a simplified LLPS partitioning mechanism (*Case 2*) in Figure 1. First, aerosol water and water-insoluble OM ($WIOM = OM - WSOC \times 1.6$) exist in two separate liquid phases, and WSOC and inorganic ions are totally dissolved in the aqueous phase. The distribution of polyol tracers between aqueous (F_w , ng m^{-3}) and WIOM (F_{WIOM} , ng m^{-3}) phases is simply depicted by their octanol-water partition coefficients (K_{OW})

$$K_{OW} = \frac{F_{WIOM}/V_{WSIOM}}{F_w/V_w} = \frac{c_{WIOM}}{c_w} \quad (4)$$

where V_{WIOM} and V_w are volumes (m^3) of WIOM and water in aerosols per cubic meter air; c_{WIOM} and c_w are solution concentrations (ng m^{-3}) of polyols concentrations in organic and aqueous phases; $\log K_{OW}$ values of target polyols were estimated using the Estimation Programs Interface (EPI) Suite developed by the US Environmental Protection Agency and Syracuse Research Corporation (Table S4; US EPA, 2012). The density of organic matter and water (ρ_w) in aerosols are set at 1.4 and 1.0 g cm^{-3} , respectively (Isaacman-VanWertz et al., 2016; Taylor et al., 2017). Second, gas-phase polyol tracers are in equilibrium with hydrophobic OM and the aqueous phase, respectively

$$K_{p,WIOM}^m = \frac{F_{WIOM}/M_{WIOM}}{A} \quad (5)$$

$$K_{H,e}^m = \frac{\frac{F_w}{M_i}}{\frac{A}{M_i \times R \times T \times \frac{c_{ALW}}{\rho_w}}} = \frac{\rho_w \times F_w}{A \times R \times T \times c_{ALW}} \quad (6)$$

where $K_{H,e}^m$ ($\text{mol m}^{-3} \text{atm}^{-1}$) is the measurement-based effective Henry's law coefficient; M_{WIOM} represents the mass concentration ($\mu\text{g m}^{-3}$) of WIOM; M_i (g mol^{-1}) is the molecular weight of specific compound; c_{ALW} ($\mu\text{g m}^{-3}$) is the mass concentration of aerosol liquid water predicted using ISORROPIA II model. *Case 3* is generally the same as *Case 2*, and the only difference is that water-soluble OM (WSOM) and WIOM exist in a single organic phase. Here total particulate OM was used instead of WIOM to assess the distribution of polyol tracers between aqueous and organic phases, and calculate partitioning coefficients of gas vs. particulate organic ($K_{p,OM}^m$) and aqueous ($K_{H,e}^m$) phases. Note that the polarity of particulate OM phase in *Case 3* was expected to increase, then using K_{OW} to calculate the distribution of polyols between organic and aqueous phases might lead to underestimated $K_{p,OM}^m$ and overestimated $K_{H,e}^m$. For comparison purposes, the Henry's law coefficient in pure water at 25 °C ($K_{H,w}^*$) was estimated using EPI and SPARC (Hilal et al., 2008; <http://archemcalc.com/sparc-web/calc>), respectively (Table S4), and was adjusted for each sampling interval due to the changes in ambient temperature using van 't Hoff equation (Text S4). ”

2. A critical question in this work, I think, what is the uncertainty on F%? Uncertainty on these measurements is not really discussed. This parameter is calculated as the ratio of two measurements, each of which likely have at least 10-15% uncertainty (typical of GC), so there is some uncertainty on F% for any given point (though that may decrease as you get to the extreme cases of being mostly in the particle phase as in this case). That doesn't account for the uncertainty on the breakthrough which is significant (e.g., methyltetrols breakthrough is $\sim 20 \pm 10\%$, so the correction factor for breakthrough is between a factor of roughly 1.1 and 1.3). Random error on each point should be reduced in the average (i.e., the average F% is known better than any one point), but the averages could be susceptible to systemic errors like uncertainty in

breakthrough that could create bias. I do notice that during the periods of high particle concentrations for e.g., 2-MTs, F% does sometimes reach 100%, so perhaps bias is minimal, but it would be nice to get some estimates of uncertainties, or additional discussion and analysis of potential error. For instance, couldn't the divergence in Figure 2 of the isoprene tracers from the line be due to some bias like uncertainty in breakthrough?

Response:

Thanks. The measurement uncertainties will not only impact the calculation results of particle phase fractions ($F\%$) of polyol tracers, but also the partitioning coefficients of gas vs. particulate organic ($K_{p,OM}^m$) and aqueous ($K_{H,e}^m$) phases.

The uncertainty estimation for measurements and partitioning coefficients were added in the revised manuscript and supplementary information, respectively.

Page 13, lines 296 – 303 of the main text:

“Uncertainty estimation. To obtain the uncertainty associated with the calculation of $F\%$ and partitioning coefficients ($K_{p,OM}^m$ and $K_{H,e}^m$), measurement uncertainties of polyol tracers in filter and PUF samples were estimated from their recoveries and breakthrough for gaseous sampling. The root sum of squares (RSS) method was applied to propagate uncertainties of gas and particle-phase concentrations for $F\%$, $K_{p,OM}^m$, and $K_{H,e}^m$ calculations. Details of the uncertainty estimation and propagation methods were provided in Text S5, and the average relative uncertainties were summarized in Table S5.”

Text S5 in supplementary information:

“Text S5. Uncertainty estimation methods

In this work, the measurement results of some polyol tracers in filter and PUF samples are subject to substantial uncertainties due to their low and variable recoveries (Table S2) and excessive breakthrough (Figure S2). A general equation was derived to estimate measurement uncertainties of individual polyols in filter and PUF samples

$$\Delta C = \sqrt{(\text{error fraction} \times \text{concentration})^2 + (0.5 \times \text{detection limit})^2} \quad (5)$$

where ΔC is the uncertainty of target species in filter (ΔQ_f and ΔQ_b , ng m^{-3}) or PUF (ΔPUF , ng m^{-3}) samples. The error fraction (%) of filter sample analysis was defined as half of the difference between maximum and minimum recoveries scaled by the average (Table S2), which was divided by (1 - average breakthrough) for PUF analysis (Figure S2). The average breakthrough of meso-erythritol (23.8%), mannose (38.1%), xylitol (36.4%), and arabitol (36.4%) were set as those of C5-alkene triols, glucose, and mannitol, respectively. According to the gas-particle separation method in this work, ΔQ_f was used to represent the uncertainty of particle-phase concentration (ΔF , ng m^{-3}), and the uncertainty of gas-phase concentration (ΔA , ng m^{-3}) was propagated by

$$\Delta A = \sqrt{\Delta Q_b^2 + \Delta \text{PUF}^2} \quad (6)$$

Then the uncertainty of total concentration (ΔS , ng m^{-3}) was calculated as

$$\Delta S = \sqrt{\Delta F^2 + \Delta A^2} \quad (7)$$

The uncertainties of particle-phase fractions ($\Delta F\%$) and partitioning coefficients ($K_{p,OM}^m$ and $K_{p,WIOM}^m$, $\text{m}^3 \text{ug}^{-1}$; $K_{H,e}^m$, $\text{mol m}^{-3} \text{atm}^{-1}$) were estimated by propagating

ΔF , ΔS , and ΔA using a simplified root sum of squares (RSS) method (Dutton et al., 2009)

$$\Delta F\% = \sqrt{\left(\frac{\partial F\%}{\partial F} \Delta F\right)^2 + \left(\frac{\partial F\%}{\partial S} \Delta S\right)^2} \times 100\% \quad (8)$$

$$\Delta K = \sqrt{\left(\frac{\partial K}{\partial F'} \Delta F'\right)^2 + \left(\frac{\partial K}{\partial A} \Delta A\right)^2} \quad (9)$$

where ΔK is the uncertainty of $K_{p,OM}^m$, $K_{p,WIOM}^m$, or $K_{H,e}^m$; F' could be F , concentrations of polyols in WIOM (F_{WIOM}) or aqueous (F_w) phases, depending on the partitioning scheme (*Cases 1–3*) and partitioning coefficient for calculation. ΔF was split into ΔF_w and ΔF_{WIOM} (or ΔF_{OM}) based on their ratios in eq. 4 of the main text. In Table S5, the estimated uncertainties are summarized and expressed in average ratios. As $K_{p,OM}^m$ and $K_{H,e}^m$ are all directly related to the ratio of particle- (F , ng m^{-3}) and gas-phase (A , ng m^{-3}) concentrations (eqs. 2, 4, 5, and 6 in the main text), their average $\Delta K/K$ values are the same (Table S5)."

Moreover, Figure 2 (now Figure 3 in the revised manuscript, shown below) has been changed by including the uncertainty of $F\%$ and one standard deviation of $\log p^{0,*}_L$ derived from different estimation tools (Table S4). Because monosaccharides (e.g., fructose) and sugar alcohols (e.g., mannitol) had low and variable recoveries (Table S2) and excessive breakthrough for gaseous sampling (Figure S2), their average $F\%$ uncertainties (6.16–31.2%) are much larger than those of isoprene SOA tracers and levoglucosan (3.33–7.24%). As shown in the figure below, more than 95% of polyols with extremely low vapor pressures ($< \sim 10^{-10}$ atm) are distributed into the particle phase, so their prominent uncertainties in $F\%$ are not physically meaningful. Considering the uncertainties in $F\%$ and $\log p^{0,*}_L$ and high average $F\%$ ($> 85\%$) of polyol tracers, a dependence of $F\%$ on the vapor pressure could not be determined.

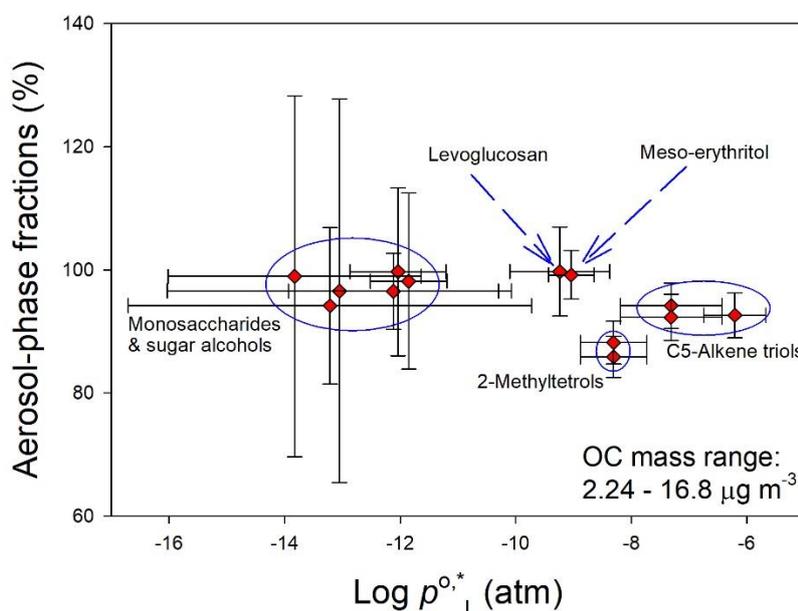


Figure 3. Average particle-phase fractions and $\log p^{0,*}_L$ of individual polyol tracers. Whiskers represent uncertainties of $F\%$ and one standard deviation of $\log p^{0,*}_L$ derived from different estimation tools.

In the revised manuscript, we added the above discussions on uncertainties of $F\%$ (Pages 17–18, lines 410–419).

“In Figure 3, The average $F\%$ uncertainties (6.16–31.2%) of monosaccharides (e.g., fructose) and sugar alcohols (e.g., mannitol) were larger than those of isoprene SOA tracers and levoglucosan (3.33–7.24%) due to their low and variable recoveries (Table S2) and excessive breakthrough (Figure S2). However, the estimated uncertainties of $F\%$ for less volatile polyols ($p^{0,*}_L < \sim 10^{-10}$ atm) were not physically meaningful, as more than 95% of these compounds existed in the particle phase. Considering the uncertainties in $F\%$ and $\log p^{0,*}_L$ and high average $F\%$ ($> 85\%$) of target polyol tracers, a dependence of $F\%$ on the vapor pressure could not be determined, and the seasonality and day-night difference ($p > 0.05$) of $F\%$ were obscured.”

3. Some of these results might be impacted by uncertainty in theoretical partitioning coefficients and by poorly constrained empirically determined coefficients.

All of the tracers shown here have values close to 100%, so there is not much dynamic range in the data and it might be susceptible to biases.

Calculation of partitioning coefficients in particular could be sensitive to uncertainties because of this (as you approach 100% in one phase, small changes in partitioning might imply large changes in K). For this reason, I'm not sure the discussion of comparison between theoretical and measured K is always that meaningful. As an extreme example, levoglucosan looks like it is always at $\sim 100\%$ except for maybe one point. Under these conditions, how can any meaningful K be measured, since a partitioning coefficient of 10^3 and 10^{100} would both produce the same effect?

On the other hand, there is substantial error in the theoretical values as well, with uncertainty in vapor pressure likely on the order of 1-2 orders of magnitude for most of these compounds (and some evidence that EPI has a tendency to overestimate compared, see Barley and McFiggans 2010).

While I agree that the time-dependent comparisons between measured and expected K against things like sulfate (e.g., Figure 3) provide insight, comments like that on line 345 comparing measured to theoretical K quantitatively aren't that meaningful.

Similarly, if you account for these sources of uncertainty, it's not clear to me that the lines in Figure 3 necessarily have a negative intercept as described. Some discussion of these uncertainties and biases might help clarify what we do know (e.g., these tracers are mostly particle phase, theoretical vapor pressures are wrong, and correlations with absorptive theory are poor), with the quantitative aspects we don't know as well (e.g., how wrong are the vapor pressures, how strong is the dependence on salt).

Response:

In our replies to the previous two comments, we have clarified the three partitioning cases defined in this work and uncertainty estimation methods for measurements and measurement-based partitioning coefficients. To address the uncertainty of vapor pressure estimated using EPI, a variety of estimation tools were deployed to provide a reasonable vapor pressure range for each polyol tracer. Referring to Table S4 of the revised supplementary information, we can find that the vapor pressure ranges of isoprene SOA tracers, levoglucosan, and meso-erythritol are within two orders of magnitude, while those for monosaccharides and mannitol are larger ($> 10^3$).

The target for comparing measured and predicted partitioning coefficients is to demonstrate that particulate OM is not the only absorbing phase, and the aerosol liquid water also plays a significant role in influencing the gas-particle partitioning of polyol tracers. In Table S5 of the revised supplementary information, the average relative uncertainties of measurement-based partitioning coefficients range from a few percent to ~50%, which will result in an uncertainty of less than ± 0.3 for logarithms of partitioning coefficients. In revised Table 1, the measurement-based $K_{p,OM}$ ($K_{p,OM}^m$) of isoprene SOA tracers, levoglucosan, and meso-erythritol for *Case 1* were more than 10 times greater than most theoretical predictions ($K_{p,OM}^t$). When the dissolution in aerosol liquid water was considered, their average $\log K_{p,OM}^m$ became much closer to or even lay within the range (e.g., levoglucosan) of $\log K_{p,OM}^t$, whenever water-soluble and water-insoluble organic matter (OM) partitioned into separate (*Cases 2*) or single (*Case 3*) liquid phases. These results support that the partitioning between gas and aerosol liquid water should not be ignored for water-soluble organic compounds like polyol tracers.

In other words, if particulate OM is the only absorbing phase in aerosols, $\log K_{p,OM}^m$ values of *Case 1* will be closer to $\log K_{p,OM}^t$ ranges, and the expected *F%* of isoprene SOA tracers, levoglucosan, and meso-erythritol should be much lower than measured in this work. To make the comparison between measured and predicted $K_{p,OM}$ meaningful, we predefined three partitioning cases and estimate vapor pressures of polyol tracers using several tools, and compared average $\log K_{p,OM}^m$ of different partitioning cases with $\log K_{p,OM}^t$ ranges. Moreover, the $K_{p,OM}^m$ uncertainty derived from measurements was estimated and would not impact the main conclusion.

The negative intercepts shown in Figures 4 and S5 of the revised manuscript and supplements are not likely ascribed to the small relative uncertainties of partitioning coefficients (Table S5). Shen et al. (2018) also identified negative intercepts of linear regressions between $\log (K_{H,w}/K_{H,e})$ and c_{sulfate} for glyoxal and methylglyoxal in the ambient atmosphere, and attributed this to unknown gas-particle partitioning mechanisms. By performing chamber experiments and comparing to existing laboratory studies, Kampf et al. (2013) found that the exponential increase of gas vs. aqueous phase partitioning coefficient ($K_{H,e}$, $\text{mol m}^{-3} \text{atm}^{-1}$) with sulfate concentration only occurred at $c_{\text{sulfate}} < 12 \text{ mol kg}^{-1} \text{ALWC}$. In Figures 4 and S5, the $\log (K_{H,w}/K_{H,e})$ values of most polyols increase faster as c_{sulfate} approaches 0, supporting that the enhanced uptake at low sulfate concentrations could be partly parameterized by the equation defining “salting-in” effects. However, the “salting-in” effect is a known phenomenon that is not likely linked with a specific physical or chemical mechanism. In previous studies, the increased partitioning of polar organic compounds to the particle phase was often attributed to organic-inorganic interactions, including reactive uptake, aqueous phase chemistry, etc.

In addition to the changes mentioned in previous two responses, we reorganized and rewrote the discussions on partitioning coefficients of gas versus organic and aqueous phases (*Sections 3.4 and 3.5*), considering the uncertainties in both measurements and predictions.

Pages 18-24, lines 420-560

“3.4 Partitioning coefficients of gas versus organic phases

To understand if particulate OM is the only absorbing phase in aerosols for polyol tracers in Nanjing, the absorptive partitioning coefficients of gas vs. organic phases were calculated based on measurement results ($K_{p,OM}^m$) for predefined *Cases*

1-3 and predicted theoretically ($K_{p,OM}^t$) using eq. 3 and vapor pressures listed in Table S4. In Table 1, $K_{p,OM}^t$ ranges of isoprene SOA tracers, levoglucosan, and meso-erythritol are within two orders of magnitude, while those of monosaccharides and mannitol are larger ($> 10^3$). When particulate OM was assumed as the only absorbing phase (*Case 1*), the average $K_{p,OM}^m$ of isoprene SOA tracers, levoglucosan, and meso-erythritol were more than 10 times greater than most of their $K_{p,OM}^t$ (Table 1), and this difference was not likely susceptible to measurement uncertainties. As shown in Table S5, the average relative uncertainties of measurement-based partitioning coefficients are all $< 50\%$, leading to an uncertainty of $\log K_{p,OM}^m$ less than ± 0.30 . Comparable or even greater (up to 10^5) gap between $K_{p,OM}^m$ and $K_{p,OM}^t$ has been observed for carbonyls in a number of laboratory and field studies (Healy et al., 2008; Zhao et al., 2013; Shen et al., 2018), which could be ascribed to reactive uptake (e.g., hydration, oligomerization, and esterification) of organic gases onto condensed phase (Galloway et al., 2009). Oligomers, sulfate and nitrate esters of 2-methyltetrols can be formed in the aerosol phase (Surratt et al., 2010; Lin et al., 2014), and their decomposition and hydrolysis during filter analysis will lead to an overestimation of particle-phase concentrations (Lin et al., 2013; Cui et al., 2018). However, the occurrence of oligomers, sulfate or nitrate esters of levoglucosan was not ever reported in ambient aerosols, although it can be readily oxidized by $\bullet OH$ in the aqueous phase of atmospheric particles (Hennigan et al., 2010; Hoffmann et al., 2010).

When solubility in aerosol liquid water was considered by assuming a LLPS in ambient aerosols, and whenever WSOM and WIOM partitioned into separate (*Case 2*) or single (*Case 3*) liquid phases, the average $\log K_{p,OM}^m$ of the above mentioned compounds became much closer to or even lay within the range (e.g., levoglucosan) of $\log K_{p,OM}^t$ (Table 1). These results indicated that the aerosol liquid water ($21.3 \pm 24.2 \mu g m^{-3}$; Table S1) is also an important absorbing phase of ambient polyol tracers in Nanjing. Similarly, the measured average $F\%$ of isoprene SOA tracers in southeastern US and central Amazonia were higher than predictions by assuming instantaneous equilibrium between the gas phase and particulate OM only, and the agreement was improved when parameterization of solubility was included for predictions (Isaacman-VanWertz et al., 2016). But none of these two studies could reasonably predict the temporal variability of $F\%$ or $\log K_{p,OM}^m$. One possible explanation is that the activity coefficients of isoprene SOA tracers and levoglucosan deviate from unity (0.42–2.04; Pye et al., 2018) and vary with PM composition. Pye et al. (2018) re-analyzed the measurement data from Isaacman-VanWertz et al. (2016) using a thermodynamic equilibrium gas-particle partitioning model in two LLPS modes, which involved organic-inorganic interactions and estimations of activity coefficients as a function of liquid PM mixture composition. The resulting predictions captured both the average and diurnal variations of measured $F\%$ for polyol tracers, suggesting a necessity in obtaining time-resolved activity coefficients for the implementation of absorptive equilibrium partitioning model.

3.5 Partitioning coefficients of gas versus aqueous phases

The predicted Henry's law coefficients in pure water ($K_{H,w}^t$, $mol m^{-3} atm^{-1}$) from EPI and SPARC estimates differed by several orders of magnitude, but literature values of isoprene SOA and levoglucosan were closer to EPI estimates (Table S4). If SPARC $K_{H,w}^*$ values were used, the average $\log K_{H,e}^m$ of most polyol tracers would be lower than their average $\log K_{H,w}^t$ (Table 2), indicating that the aqueous phase of ambient aerosol is less hospitable to polyol tracers than pure water. This is in conflict with the fact that the interactions of organic compounds, water, and inorganic ions in

aerosols will increase the partitioning of highly oxygenated compounds ($O:C \geq 0.6$; e.g., isoprene SOA tracers and levoglucosan) into the particle phase (Pye et al., 2018). Several studies identified a close relationship between salt concentrations of aerosol water and enhanced uptake of very polar compounds (Almeida et al., 1983; Kroll et al., 2005; Ip et al., 2009; Kampf et al., 2013; Shen et al., 2018). Thus, $\log K_{H,w}^t$ values of EPI estimates were used for further data analysis.

In Table 2, the $K_{H,w}^t$ values of isoprene SOA tracers, levoglucosan, and meso-erythritols based on EPI estimations were 10^2 to 10^6 lower than their corresponding $K_{H,e}^m$. $\log K_{H,e}^m$ values of *Cases 2* and *3* had ignorable difference and were not presented separately. Other polyol compounds exhibited less difference between $\log K_{H,e}^m$ and $\log K_{H,w}^t$, which was very likely caused by the overestimation of their gas-phase concentrations. The average $K_{H,e}^m$ values of polyol tracers (10^{13} – 10^{15} mol m^{-3} atm $^{-1}$) in this study were several orders of magnitude larger than those of carbonyls derived from ambient measurements (10^{10} – 10^{12} mol m^{-3} atm $^{-1}$; Shen et al., 2018) and chamber simulations ($\sim 10^{11}$ mol m^{-3} atm $^{-1}$; Kroll et al., 2005; Volkamer et al., 2006; Galloway et al., 2009). This is because low molecular weight carbonyls (e.g., glyoxal) are much more volatile ($p^{\circ,*}_L > 10^{-2}$ atm) than our target polyols (Table S4). According to existing studies, the minimum concentrations of gas-phase glyoxal and methylglyoxal in Chinese cities (~ 0.1 $\mu\text{g m}^{-3}$) are magnitudes higher than the averages of polyol tracers in this work, while their particle-phase concentrations are of the same magnitude (Shen et al., 2018; Liu et al., 2020).

A number of previous studies observed enhanced $K_{H,e}$ of carbonyls with salt concentrations in aqueous solution (Ip et al., 2009; Kampf et al., 2013; Waxman et al., 2015; Shen et al., 2018), and described this “salting-in” effect using

$$\text{Log} \left(\frac{K_{H,w}}{K_{H,e}} \right) = K_s c_{\text{salt}} \quad (7)$$

where K_s (kg mol $^{-1}$) is the salting constant, and c_{salt} is the aqueous-phase concentration of salt in mol kg $^{-1}$ ALWC. This equation is originally defined in Setschenow (1889) by plotting $\log (K_{H,w}/K_{H,e})$ versus the total salt concentration (mol L $^{-1}$).

As sulfate has been identified as the major factor influencing the salting effect of carbonyl species (Kroll et al., 2005; Ip et al., 2009), Figure 4 shows modified Setschenow plots for C5-alkene triols, 2-methyltetrols, and levoglucosan, where $\log (K_{H,w}^t/K_{H,e}^m)$ values were regressed to the molality of sulfate ion in aerosol liquid water (c_{sulfate} , mol kg $^{-1}$ ALWC). The $\log (K_{H,w}^t/K_{H,e}^m)$ data increased faster when c_{sulfate} approached 0, and deviated from their expected behavior with increased c_{sulfate} . Kampf et al. (2013) selected a threshold c_{sulfate} of 12 mol kg $^{-1}$ ALWC to illustrate the deviation for chamber experiments, and attributed it to elevated viscosity and slow particle-phase reactions at high c_{sulfate} . In Figure 4, negative correlations ($p < 0.01$) are observed at $c_{\text{sulfate}} < 12$ mol kg $^{-1}$ ALWC, and Figure S5 exhibits significant negative correlations ($p < 0.01$) between $\log (K_{H,w}^t/K_{H,e}^m)$ and c_{sulfate} for individual polyols even without excluding the deviations at high c_{sulfate} . The K_s values of polyol tracers from Figures 4 and S5 (-0.17 – -0.037 kg mol $^{-1}$) are in a similar range as that of glyoxal (-0.24 – -0.04 kg mol $^{-1}$; Kampf et al., 2013; Shen et al., 2018; Waxman et al., 2015). These results indicated that the shifting of gas-particle equilibrium toward the condensed phase might be partly parameterized by the equation defining “salting-in” effects.

However, the “salting-in” effect is a known phenomenon that is not likely linked with a specific physical or chemical mechanism. Quantum chemical calculation results indicated negative Gibbs free energy of water displacement for interactions between SO_4^{2-} and glyoxal monohydrate (Waxman et al., 2015). The net “salting-in”

effect of 1-nitro-2-naphthol in NaF solution was interpreted by postulating hydrogen bonding (Almeida et al., 1983). A direct binding of cations to ether oxygens was proposed to be responsible for the increased solubility of water-soluble polymers (Sadeghi and Jahani, 2012). Due to the moderate correlations and negative intercepts in Figures 4 and S5, the gap between $K_{H,e}^l$ and $K_{H,w}^m$ cannot be closed by the “salting-in” effect alone. Shen et al. (2018) also obtained negative intercepts when plotting $\log(K_{H,w}^l/K_{H,e}^m)$ over c_{sulfate} for glyoxal and methylglyoxal in ambient atmosphere, and attributed this to unknown gas-particle partitioning mechanisms. Evidences showing that the thermal degradation of less volatile oligomers and organosulfates can lead to an overestimation of 2-methyltetrols by 60–188% when using a conventional GC/EI-MS method (Cui et al., 2018). To fit the gas-particle distribution of 2-methyltetrols in southeastern US, 50% of particulate 2-methyltetrols was presumed to exist in chemical forms with much lower vapor pressures by Pye et al. (2018). So, the reactive uptake and aqueous phase chemistry could be explanations for the enhanced uptake of isoprene SOA tracers. Moreover, $\log(K_{H,w}^l/K_{H,e}^m)$ values of polyol tracers also negatively correlated with the aqueous-phase concentrations of WSOC (c_{WSOC} , Figure S6), but not NH_4^+ or NO_3^- . This dependence might be associated with the “like-dissolves-like” rule, or indicate the importance of aqueous-phase heterogeneous reactions (Hennigan et al., 2009; Volkamer et al., 2009). Although several studies have estimated Henry’s law constants for a variety of polar organic compounds in pure water (e.g., polyols and polyacids; Compornolle and Müller, 2014a, b), more work is warranted to decrease the estimation uncertainty and explain their increased partitioning toward aerosol liquid water explicitly.”

Specific comments:

4. Line 57: Typo, "Filed" should say "Field"

Response:

It has been revised as suggested. (Page 4, line 80)

5. Line 85: Probably also worth mentioning that Yatavelli et al. (2014, 10.5194/acp-14-1527-2014) and Isaacman-VanWertz et al. (2016) also found good agreement with theory for alkanolic acids

Response:

The two references were added, and the original expression has been changed into

“Unlike non-polar species (e.g., *n*-alkanes, polycyclic aromatic hydrocarbons) and alkanolic acids that are well simulated (Simcik et al., 1998; Xie et al., 2014a; Yatavelli et al., 2014; Isaacman-VanWertz et al., 2016), particle-phase concentrations of carbonyls were underestimated by several orders of magnitude when particulate OM is presumed as the only absorbing phase in ideal condition (Healy et al., 2008; Kampf et al., 2013; Shen et al., 2018).” (Page 5, 109-114)

6. Line 90: Typo, "every" should say "very"

Response:

It has been revised as suggested. (Page 5, line 117)

7. Line 176: When quantifying the isoprene tracers using meso-erythritol, I assume the effect of fragmentation on the quant ion was accounted for, but that is not clear here or in the supplemental. In other words, while m/z 219 is used for the 2-MTs, something different, likely m/z 217, is used for meso-erythritol, and the fraction of total signal that is the quant ion could be different - if this was corrected for it should be stated at least in the SI, if not it should be justified.

Response:

In the current work, we did not consider the effect of fragmentation on the quant ion due to the following reasons:

(1) Meso-erythritol and isoprene tracers (C5-alkene triols and 2-methyltetrols) should have different total ion signal intensity for the same amount, so the calibration curve of meso-erythritol is expected to differ from isoprene tracers. Even if the effect of fragmentation on the quant ion was accounted for, the quantification accuracy might not be improved. A number of previous studies used meso-erythritol as surrogate for isoprene tracers without considering the effect of fragmentation (Claeys et al., 2004; Hu et al., 2008; Ding et al., 2008, 2012; Lin et al., 2013; Xie et al., 2014). It should also be noted that ketopenic acid was used to quantify all SOA tracers by Kleindienst et al. (2007).

(2) The present study focused on the gas-particle partitioning of polyol tracers. Their particle-phase fractions, measurement-based partitioning coefficients of gas versus organic ($K_{p,OM}^m$) and aqueous ($K_{H,e}^m$) phases were all directly related to the ratio of particle- to gas-phase concentrations, which is not impacted by the effect of fragmentation.

(3) The total ion peaks corresponding to isoprene tracers in ambient air samples usually combined with other compounds, thus the quant ions (m/z 219) fractions of 2-methyltetrols in total ion signal (0.14 ± 0.035 , 0.20 ± 0.015) had larger variability than that of meso-erythritol (0.12 ± 0.0039), then the correction for fragmentation effect might introduce new uncertainty.

In the revised supplementary information, we added some justifications for not correcting the fragmentation effect.

“Meso-erythritol and isoprene SOA tracers were expected to have different total ion intensity for the same amount, and the total ion signals of isoprene tracers in ambient air sample often co-eluted with other compounds, so the fragmentation difference between quantification ions of meso-erythritol (m/z 217) and isoprene SOA tracers (m/z 219) was not adjusted in this work. It should be noted that meso-erythritol was used as the surrogate for isoprene tracers without considering the fragmentation effect in a number of previous studies (Claeys et al., 2004; Hu et al., 2008; Ding et al., 2008, 2012; Lin et al., 2013; Xie et al., 2014).”

8. Line 197: If gas-phase mass is being taken as Q_b plus PUF, what is the purpose of the Q_b measurement at all? Why not just do Q_f backed by a PUF?

Response:

The backup quartz fiber filter was typically used to assess positive sampling artifact of particulate OC due to gaseous adsorption (Chow et al., 2010; Subramanian et al., 2004). In a companion study to this work, concentrations of bulk species in $PM_{2.5}$ were determined by subtracting measurement results of Q_b from Q_f (Yang et al., 2021). This has been mentioned in section 2.2 of the manuscript (Pages 8, lines 176-178).

Concentrations of polyol tracers detected on Q_b reflected filter-based sampling artifacts, which is ascribed to gaseous adsorption (“positive artifact”) and evaporation loss from particles on Q_f (“negative artifact”). If polyol tracers on Q_b account for significant fractions of their total concentrations, then the separation of their particle- and gas-phase concentrations will depend largely on sampling artifact corrections. In the current work, polyol tracers were predominantly observed on Q_f with averages 1-3 orders of magnitude higher than those on Q_b and PUF, and it would be safe to assume that Q_b values were equally contributed by positive and negative artifacts. So, Q_b measurement is necessary in determining the magnitude of sampling artifacts.

The original expressions have been revised (Pages 9-10, lines 218-229).

“Polyol tracers detected in Q_b samples are contributed by both gaseous adsorption (“positive artifact”) and particle-phase evaporation from Q_f samples (“negative artifact”), but their relative contributions are unknown. Xie et al. (2014b) adjusted particle- and gas-phase concentrations of levoglucosan and 2-methyltetrol based on Q_b measurements in two different ways. One assumed that Q_b values were completely attributed to gaseous adsorption; the other presumed equal contributions from gaseous adsorption and Q_f evaporation. However, negligible difference in gas-particle distribution was observed due to the small Q_b values. In Table S3, concentrations of polyol tracers on Q_b are far below those on Q_f , and it would be safe to presume equal positive and negative artifacts. In this study, particle-phase concentrations of polyols were represented by Q_f values, and the gas phase was calculated as the sum of Q_b and PUF measurements.”

9. Line 220-225: I'm not completely clear on the partitioning approach. The K_{ow} is used to partition the particle fraction between condensed phases - is this information then used in the gas-particle partitioning? For instance, is K_{OM} used for the organic component, and K_H used for the aqueous? I guess I'm just not clear on how K_{ow} fits into the scheme.

Response:

Yes, the K_{ow} is used to calculate concentrations of polyol tracers in two separate aerosol phases (organic and aqueous). $K_{p,OM}$ and K_H represent gas-particle partitioning coefficients of gas versus organic and aqueous phases, respectively.

To make it clear, we defined three partitioning cases, and elucidate the calculation methods for the two partitioning coefficients in the revised manuscript.

Pages 10-13, lines 238-295 (See responses to *Comment 1*, eqs 4, 5, and 6)

10. Line 297-298: Isaacman-VanWertz et al. (2016) also show hourly diurnal profiles with a daytime high, it is interesting to see in this work that the difference between daytime and nighttime values was not significant. As noted later in the manuscript, in the summer when there are actually strong concentrations of these compounds, there is a strong diurnal, so I'm not sure it really makes sense to claim here there is no diurnal.

Response:

Similar to the observations by Fu and Kawamura (2011) at a forest site in Hokkaido, Japan, concentrations of isoprene SOA tracers in southeastern US and central Amazonia also exhibited peak concentrations from late afternoon to mid night. In the current work, daytime and nighttime samples were collected during 8:00 AM – 7:00 PM and 7:00 PM – 7:00 AM (next day), respectively. Neither the daytime or nighttime sample covered the whole period when isoprene SOA had peak concentrations. This explains why the strong diurnal variations of isoprene SOA tracers were not captured in this work.

The original expression was changed into

“Previous field studies observed strong diurnal variations of isoprene SOA tracers with peak concentrations from afternoon till midnight (Fu and Kawamura, 2011; Isaacman-VanWertz et al., 2016). Although no IEPOX will be generated from the oxidation of isoprene by $\bullet\text{OH}$ and $\text{HO}_2\bullet$ after sunset, the formations of C5-alkene triols and 2-methyltetrols might continue until pre-existing IEPOX is exhausted. In this work, neither the daytime (8:00 AM–7:00 PM) or night-time (7:00 AM–7:00 AM next day) sample covered the whole period when isoprene SOA tracers had peak concentrations, and the strong diurnal variations of C5-alkene triols and 2-methyltetrols were not captured.” (Page 16, 367-375)

11. Line 336: This is the first place $F\%$ is given any error or range, though I assume here the uncertainty is the standard deviation.

Response:

We clarified in the revised manuscript that the numbers here represented ranges of average \pm standard deviation. (Page 17, lines 405)

Moreover, measurement uncertainties and their influences on $F\%$ and partitioning coefficients were estimated in the revised manuscript. (Text S5 and Figure 3, see responses to *Comment 2*; Table S5, shown below)

Table S5. Average relative uncertainties of measurements and calculated parameters.

Species	$\Delta F/F^a$	$\Delta A/A^b$	$\Delta S/S^c$	$\Delta F\%/F\%^d$	$\Delta K/K^e$
Isoprene SOA tracers					
C5-alkene triol 1	0.028	0.032	0.027	0.037	0.043
C5-alkene triol 2	0.028	0.054	0.033	0.036	0.059
C5-alkene triol 3	0.028	0.077	0.034	0.038	0.084
2-Methylthreitol	0.028	0.051	0.028	0.033	0.059
2-Methylerythritol	0.028	0.066	0.030	0.035	0.072
Biomass burning tracer					
Levoglucosan	0.051	0.16	0.054	0.072	0.17
Sugars and sugar alcohols					
Meso-erythritol	0.028	0.11	0.028	0.040	0.12
Fructose	0.23	0.27	0.26	0.31	0.36
Mannose	0.045	0.27	0.049	0.062	0.28
Glucose	0.094	0.28	0.10	0.18	0.31
Xylitol	0.10	0.12	0.10	0.14	0.16
Arabitol	0.097	0.26	0.099	0.14	0.28
Mannitol	0.21	0.42	0.21	0.29	0.47

^a Particle-phase concentration; ^b gas-phase concentration; ^c total concentration; ^d particle-phase fraction; ^e partitioning coefficients of gas vs. organic and aqueous phases.

12. Line 370: This could also be due to some systematic bias in how EPI estimates vapor pressures. This might be reconciled by testing other vapor pressure estimation methods. For example EVAPORATION agrees with EPI on the vapor pressure of erythritol ($\log v_p$ (atm) = $\sim -8.2-8.5$) but EPI estimates mannose to be two orders of magnitude more volatile ($\log v_p$ (atm) = -9.5) than EVAPORATION estimates ($\log v_p$ (atm) = -11.5). This also highlights the uncertainty of using comparisons against theoretical K to draw conclusions.

Response:

To address the concern on uncertainties of vapor pressures, other estimation tools were tested (Table S4). The revised Table 1 (shown below) compared measurement-based $\log K_{p,OM}$ of predefined Cases 1-3 and theoretical predictions based on different vapor pressure estimation methods. Then the discussions on comparisons between measured and predicted $K_{p,OM}$ were revised accordingly (Section 3.4, see responses to Comment 3; pages 18-20, lines 420-480).

As shown in Table 1, when the solubility in aerosol liquid water is considered (Cases 2 and 3), the agreement between $\log K_{p,OM}^m$ and $\log K_{p,OM}^t$ for isoprene SOA tracers and levoglucosan has been improved substantially. While the deviations of $\log K_{p,OM}^m$ versus $\log K_{p,OM}^t$ for monosaccharides and sugar alcohols became larger. This is expected to be caused by the overestimations of their gas phase concentrations due to sampling artifacts, since the vapor pressures of sugar polyols were orders of magnitude lower than isoprene SOA tracers and levoglucosan.

Table 1. Comparisons of measurement-based $\log K_{p,OM}$ ($m^3 \mu g^{-1}$) at three proposed cases and predicted values.

Species	No. of obs.	Log $K_{p,OM}^m$ ^a			Log $K_{p,OM}^t$ ^b			
		Case 1	Case 2	Case 3	EPI	EVAPORATION	SPARC	SIMPOL
Isoprene SOA tracers								
C5-alkene triol 1	53	0.33 ± 0.71	-0.79 ± 0.86	-0.82 ± 0.85	-3.09	-2.84	-1.19	-2.88
C5-alkene triol 2	63	0.15 ± 0.55	-1.02 ± 0.74	-1.05 ± 0.73	-3.62	-3.67	-4.14	-2.85
C5-alkene triol 3	83	0.35 ± 0.68	-0.83 ± 0.86	-0.86 ± 0.85	-2.90	-2.65	-1.00	-2.69
2-Methylthreitol	101	-0.12 ± 0.48	-2.09 ± 0.71	-2.09 ± 0.70	-1.87	-1.30	-1.18	-0.47
2-Methylerythritol	95	-0.011 ± 0.58	-1.96 ± 0.71	-1.96 ± 0.71	-1.90	-1.34	-1.22	-0.50
Biomass burning tracer								
Levoglucofan	65	2.23 ± 0.72	0.63 ± 0.90	0.62 ± 0.90	-0.04	-0.81	1.04	-0.76
Sugars and sugar alcohols								
Meso-erythritol	31	0.87 ± 0.53	-1.43 ± 0.60	-1.43 ± 0.60	-0.65	-1.21	-0.45	
Fructose	85	0.65 ± 0.73	-1.20 ± 0.83	-1.20 ± 0.89	1.17	2.76	6.94	
Mannose	74	0.62 ± 0.71	-2.12 ± 0.95	-2.12 ± 0.95	1.28	2.13	4.77	
Glucose	88	0.42 ± 0.67	-2.77 ± 0.93	-2.77 ± 0.93	0.34	3.75	7.32	
Xylitol	22	0.24 ± 0.54	-2.61 ± 0.72	-2.61 ± 0.72	3.37	2.34	3.57	
Arabitol	30	1.46 ± 0.89	-1.35 ± 1.24	-1.35 ± 1.24	3.25	1.67	2.90	
Manitol	65	1.08 ± 0.63	-2.24 ± 0.95	-2.24 ± 0.95	2.33	4.16	6.68	

^a Average ± standard deviation; ^b temperature range: -4–36 °C.

13. Line 391: Here and throughout, why use $K_{H,e}$ and $K_{H,w}$ for measured and predicted, respectively, for Henry's law, but K_p^m and K_p^t or measured and predicted, respectively, for absorptive partitioning. I think the discussion would be more clear if notation were more consistent (e.g, K_H^m and K_H^t)

Response:

Thanks. To make it more clear and consistent, $K_{H,e}$ and $K_{H,w}$ were changed into $K_{H,e}^m$ and $K_{H,w}^t$ if necessary throughout the manuscript.

14. Line 407: out of curiosity, why did the authors choose to switch to molality instead of molarity?

Response:

The switch from molarity to molality will not influence the correlation between $\log(K_{H,w}/K_{H,e})$ and sulfate concentrations. It also makes it easier to compare the salting constant (K_s) with previous studies (Kampf et al., 2013; Waxman et al., 2015; Shen et al., 2018).

References

- Chow, J. C., Watson, J. G., Chen, L. W. A., Rice, J., and Frank, N. H.: Quantification of PM_{2.5} organic carbon sampling artifacts in US networks, *Atmos. Chem. Phys.*, 10, 5223-5239, 10.5194/acp-10-5223-2010, 2010.
- Claeys, M., Graham, B., Vas, G., Wang, W., Vermeylen, R., Pashynska, V., Cafmeyer, J., Guyon, P., Andreae, M. O., Artaxo, P., and Maenhaut, W.: Formation of secondary organic aerosols through photooxidation of isoprene, *Science*, 303, 1173-1176, 10.1126/science.1092805, 2004.
- Ding, X., Zheng, M., Yu, L., Zhang, X., Weber, R. J., Yan, B., Russell, A. G., Edgerton, E. S., and Wang, X.: Spatial and seasonal trends in biogenic secondary organic aerosol tracers and water-soluble organic carbon in the Southeastern United States, *Environ. Sci. Technol.*, 42, 5171-5176, 10.1021/es7032636, 2008.
- Ding, X., Wang, X.-M., Gao, B., Fu, X.-X., He, Q.-F., Zhao, X.-Y., Yu, J.-Z., and Zheng, M.: Tracer-based estimation of secondary organic carbon in the Pearl River Delta, south China, *J. Geophys. Res. Atmos.*, 117, D05313, 10.1029/2011jd016596, 2012.
- Hu, D., Bian, Q., Li, T. W. Y., Lau, A. K. H., and Yu, J. Z.: Contributions of isoprene, monoterpenes, β -caryophyllene, and toluene to secondary organic aerosols in Hong Kong during the summer of 2006, *J. Geophys. Res. Atmos.*, 113, D22206, 10.1029/2008jd010437, 2008.
- Kleindienst, T. E., Jaoui, M., Lewandowski, M., Offenber, J. H., Lewis, C. W., Bhave, P. V., and Edney, E. O.: Estimates of the contributions of biogenic and anthropogenic hydrocarbons to secondary organic aerosol at a southeastern US location, *Atmos. Environ.*, 41, 8288-8300, 10.1016/j.atmosenv.2007.06.045, 2007.
- Kampf, C. J., Waxman, E. M., Slowik, J. G., Dommen, J., Pfaffenberger, L., Praplan, A. P., Prévôt, A. S. H., Baltensperger, U., Hoffmann, T., and Volkamer, R.: Effective Henry's law partitioning and the salting constant of glyoxal in aerosols containing sulfate, *Environ. Sci. Technol.*, 47, 4236-4244, 10.1021/es400083d, 2013.
- Lin, Y. H., Knipping, E. M., Edgerton, E. S., Shaw, S. L., and Surratt, J. D.: Investigating the influences of SO₂ and NH₃ levels on isoprene-derived secondary organic aerosol formation using conditional sampling approaches, *Atmos. Chem. Phys.*, 13, 8457-8470, 10.5194/acp-13-8457-2013, 2013.
- Shen, H., Chen, Z., Li, H., Qian, X., Qin, X., and Shi, W.: Gas-particle partitioning of carbonyl compounds in the ambient atmosphere, *Environ. Sci. Technol.*, 52, 10997-11006, 10.1021/acs.est.8b01882, 2018.
- Subramanian, R., Khlystov, A. Y., Cabada, J. C., and Robinson, A. L.: Positive and negative artifacts in particulate organic carbon measurements with denuded and undenuded sampler configurations special issue of aerosol science and technology on findings from the fine particulate matter supersites program, *Aerosol Sci. Tech.*, 38, 27-48, 10.1080/02786820390229354, 2004.
- Waxman, E. M., Elm, J., Kurtén, T., Mikkelsen, K. V., Ziemann, P. J., and Volkamer, R.: Glyoxal and methylglyoxal Setschenow salting constants in sulfate, nitrate, and chloride solutions: Measurements and Gibbs energies, *Environ. Sci. Technol.*, 49, 11500-11508, 10.1021/acs.est.5b02782, 2015.
- Xie, M., Hannigan, M. P., and Barsanti, K. C.: Gas/particle partitioning of 2-methyltetrols and levoglucosan at an urban site in Denver, *Environ. Sci. Technol.*, 48, 2835-2842, 10.1021/es405356n, 2014.
- Yang, L., Shang, Y., Hannigan, M. P., Zhu, R., Wang, Q. g., Qin, C., and Xie, M.: Collocated speciation of PM_{2.5} using tandem quartz filters in northern nanjing, China: Sampling artifacts and measurement uncertainty, *Atmos. Environ.*, 246, 118066, <https://doi.org/10.1016/j.atmosenv.2020.118066>, 2021.

This study presents the results of simultaneous gas- and particle-phase measurements of oxygenated organic compounds in Nanjing, China. Due to some potential measurement artifacts with some compounds, the results focus on C5-alkene triols, 2-methyltetrols, and levoglucosan. The major finding is that the particle-phase fraction of these compounds were on average orders of magnitude higher than can be explained by either absorptive partitioning theory or Henry's law. There is moderate evidence that sulfate caused a "salting in" effect, though more discussion and/or data are needed to support this point (comment detailed below). The writing and organization are generally good, and the topic is of interest to a broad audience. I recommend the manuscript for publication after the following comments are addressed.

Response:

Thanks for the reviewer's comments, and we will reply these point by point in the reviewers' specific comments.

Specific Comments

1. How much does measurement uncertainty affect the partitioning coefficients? The stated acceptable threshold for breakthrough (< 33%) seems relatively high. How do the observed levels of breakthrough contribute to the uncertainty in the partitioning coefficients? Further, did breakthrough vary as a function of ambient temperature, OA loading, etc? Discussion of these points is needed.

Response:

The estimation of measurement uncertainties and their influences on partitioning coefficients were added in the revised manuscript. In general, measurement uncertainties of polyol tracers in filter samples were estimated from averages and ranges of their recoveries and method detection limits. Breakthrough of gaseous sampling was additionally considered for measurement uncertainties of PUF samples. Uncertainties associated with the calculations of particle-phase fractions and partitioning coefficients were estimated using a simplified root sum of squares (RSS) method by propagating measurement uncertainties of gas- and particle-phase concentrations. Details of the method were provided in Text S5 in supplementary information and mentioned in the main text.

Text S5 in supplementary information:

“Text S5. Uncertainty estimation methods

In this work, the measurement results of some polyol tracers in filter and PUF samples are subject to substantial uncertainties due to their low and variable recoveries (Table S2) and excessive breakthrough (Figure S2). A general equation was derived to estimate measurement uncertainties of individual polyols in filter and PUF samples

$$\Delta C = \sqrt{(\text{error fraction} \times \text{concentration})^2 + (0.5 \times \text{detection limit})^2} \quad (5)$$

where ΔC is the uncertainty of target species in filter (ΔQ_f and ΔQ_b , ng m⁻³) or PUF (ΔPUF , ng m⁻³) samples. The error fraction (%) of filter sample analysis was defined as half of the difference between maximum and minimum recoveries scaled by the average (Table S2), which was divided by (1 - average breakthrough) for PUF analysis (Figure S2). The average breakthrough of meso-erythritol (23.8%), mannose (38.1%), xylitol (36.4%), and arabitol (36.4%) were set as those of C5-alkene triols, glucose, and

mannitol, respectively. According to the gas-particle separation method in this work, ΔQ_f was used to represent the uncertainty of particle-phase concentration (ΔF , ng m⁻³), and the uncertainty of gas-phase concentration (ΔA , ng m⁻³) was propagated by

$$\Delta A = \sqrt{\Delta Q_b^2 + \Delta PUF^2} \quad (6)$$

Then the uncertainty of total concentration (ΔS , ng m⁻³) was calculated as

$$\Delta S = \sqrt{\Delta F^2 + \Delta A^2} \quad (7)$$

The uncertainties of particle-phase fractions ($\Delta F\%$) and partitioning coefficients ($K_{p,OM}^m$ and $K_{p,WIOM}^m$, m³ ug⁻¹; $K_{H,e}^m$, mol m⁻³ atm⁻¹) were estimated by propagating ΔF , ΔS , and ΔA using a simplified root sum of squares (RSS) method (Dutton et al., 2009)

$$\Delta F\% = \sqrt{\left(\frac{\partial F\%}{\partial F} \Delta F\right)^2 + \left(\frac{\partial F\%}{\partial S} \Delta S\right)^2} \times 100\% \quad (8)$$

$$\Delta K = \sqrt{\left(\frac{\partial K}{\partial F'} \Delta F'\right)^2 + \left(\frac{\partial K}{\partial A} \Delta A\right)^2} \quad (9)$$

where ΔK is the uncertainty of $K_{p,OM}^m$, $K_{p,WIOM}^m$, or $K_{H,e}^m$; F' could be F , concentrations of polyols in WIOM (F_{WIOM}) or aqueous (F_w) phases, depending on the partitioning scheme (*Cases 1–3*) and partitioning coefficient for calculation. ΔF was split into ΔF_w and ΔF_{WIOM} (or ΔF_{OM}) based on their ratios in eq. 4 of the main text. In Table S5, the estimated uncertainties are summarized and expressed in average ratios. As $K_{p,OM}^m$ and $K_{H,e}^m$ are all directly related to the ratio of particle- (F , ng m⁻³) and gas-phase (A , ng m⁻³) concentrations (eqs. 2, 4, 5, and 6 in the main text), their average $\Delta K/K$ values are the same (Table S5). ”

Page 13, lines 296 – 303 of the main text:

“**Uncertainty estimation.** To obtain the uncertainty associated with the calculation of $F\%$ and partitioning coefficients ($K_{p,OM}^m$ and $K_{H,e}^m$), measurement uncertainties of polyol tracers in filter and PUF samples were estimated from their recoveries and breakthrough for gaseous sampling. The root sum of squares (RSS) method was applied to propagate uncertainties of gas and particle-phase concentrations for $F\%$, $K_{p,OM}^m$, and $K_{H,e}^m$ calculations. Details of the uncertainty estimation and propagation methods were provided in Text S5, and the average relative uncertainties were summarized in Table S5.”

According to the equation for breakthrough calculation, a value of 33% means that the amount of a certain compound in backup PUF samples is half of that in front PUF samples. It was used as an indicator of excessive breakthrough in several previous studies (Peters et al., 2000; Ahrens et al., 2011; Xie et al., 2014a, b). However, the breakthrough value was rarely used to correct measurement results.

In the revised manuscript, the breakthrough of gaseous sampling was included for the estimation of measurement uncertainties and their influences on partitioning coefficients, but its individual contributions cannot be separated. In Table S5 (shown below), uncertainties of gas- and particle-phase concentrations, particle-phase fractions, and partitioning coefficients are expressed in average relative abundance.

Due to the limit in sample number for breakthrough tests and low detection rates, we can hardly evaluate the dependence of breakthrough on ambient temperature or OA loadings. The breakthrough of an ideal sampling method is expected to be extremely low (e.g., <10%) and have no dependence on ambient temperature or OA loadings. These statements were added in the revised manuscript (Pages 14-15, 337–341).

Table S5. Average relative uncertainties of measurements and calculated parameters.

Species	$\Delta F/F^a$	$\Delta A/A^b$	$\Delta S/S^c$	$\Delta F\%/F\%^d$	$\Delta K/K^e$
Isoprene SOA tracers					
C5-alkene triol 1	0.028	0.032	0.027	0.037	0.043
C5-alkene triol 2	0.028	0.054	0.033	0.036	0.059
C5-alkene triol 3	0.028	0.077	0.034	0.038	0.084
2-Methylthreitol	0.028	0.051	0.028	0.033	0.059
2-Methylerythritol	0.028	0.066	0.030	0.035	0.072
Biomass burning tracer					
Levoglucosan	0.051	0.16	0.054	0.072	0.17
Sugars and sugar alcohols					
Meso-erythritol	0.028	0.11	0.028	0.040	0.12
Fructose	0.23	0.27	0.26	0.31	0.36
Mannose	0.045	0.27	0.049	0.062	0.28
Glucose	0.094	0.28	0.10	0.18	0.31
Xylitol	0.10	0.12	0.10	0.14	0.16
Arabitol	0.097	0.26	0.099	0.14	0.28
Mannitol	0.21	0.42	0.21	0.29	0.47

^a Particle-phase concentration; ^b gas-phase concentration; ^c total concentration; ^d particle-phase fraction; ^e partitioning coefficients of gas vs. organic and aqueous phases.

2. The assumption of LLPS should be discussed. Other studies, for example Pye et al. (2018), could be included in this discussion.

Response:

Thanks. We discussed the assumption of LLPS by citing two existing studies (Zuend and Seinfeld, 2012; Pye et al., 2018) in the revised manuscript. Pye et al. (2018) was included in discussions on the agreement between measurement-based and predicted partitioning coefficients.

Page 11, lines 258-262

“Due to the influence of mixing state and water content in aerosols, several studies modeled the gas-particle partitioning of oxygenated organic compounds by defining a liquid-liquid phase separation (LLPS) in the aerosol (Zuend and Seinfeld, 2012; Pye et al., 2018). The organic-inorganic interactions and changes of activity coefficients in aqueous mixtures were fully considered as well.”

Pages 19-20, lines 444-464

“When solubility in aerosol liquid water was considered by assuming a LLPS in ambient aerosols, and whenever WSOM and WIOM partitioned into separate (*Case 2*) or single (*Case 3*) liquid phases, the average $\log K_{p,OM}^m$ of the above mentioned compounds became much closer to or even lay within the range (e.g., levoglucosan) of $\log K_{p,OM}^l$ (Table 1). These results indicated that the aerosol liquid water ($21.3 \pm 24.2 \mu\text{g m}^{-3}$; Table S1) is also an important absorbing phase of ambient polyol tracers in Nanjing. Similarly, the measured average $F\%$ of isoprene SOA tracers in southeastern US and central Amazonia were higher than predictions by assuming instantaneous equilibrium between the gas phase and particulate OM only, and the agreement was improved when parameterization of solubility was included for predictions (Isaacman-

VanWertz et al., 2016). But none of these two studies could reasonably predict the temporal variability of $F\%$ or $\log K_{p,OM}^m$. One possible explanation is that the activity coefficients of isoprene SOA tracers and levoglucosan deviate from unity (0.42–2.04; Pye et al., 2018) and vary with PM composition. Pye et al. (2018) re-analyzed the measurement data from Isaacman-VanWertz et al. (2016) using a thermodynamic equilibrium gas-particle partitioning model in two LLPS modes, which involved organic-inorganic interactions and estimations of activity coefficients as a function of liquid PM mixture composition. The resulting predictions captured both the average and diurnal variations of measured $F\%$ for polyol tracers, suggesting a necessity in obtaining time-resolved activity coefficients for the implementation of absorptive equilibrium partitioning model.”

Page 21, lines 487-490

“This is in conflict with the fact that the interactions of organic compounds, water, and inorganic ions in aerosols will increase the partitioning of highly oxygenated compounds ($O:C \geq 0.6$; e.g., isoprene SOA tracers and levoglucosan) into the particle phase (Pye et al., 2018).”

3. I think that the title is somewhat misleading because the answer to the question is actually “neither” for most of the organic markers investigated. I suggest revising the title to reflect this.

Response:

In the revised manuscript, it was shown that the solubility of polyol tracers in aerosol liquid water should not be ignored. Comparisons of measurement-based effective Henry’s law constants versus predicted values in pure water indicated increased partitioning toward the particle phase.

Then the title has been changed into

“Gas-particle partitioning of polyol tracers at a suburban site in Nanjing, east China: Increased partitioning to the particle phase”

4. The comparisons to Denver, CO seem completely random given that meteorology, OA loadings, inorganic composition, and ALWC are quite different between the two locations. I understand that this research group made measurements in both locations, but some additional discussion is warranted to better connect the two locations.

Response:

The observation study in Denver, CO (Xie et al., 2014b) did not measure inorganic composition or estimate ALWC. In that study, evidences showing that gas-particle partitioning of 2-methyltetrol and levoglucosan depended on variations in ambient temperature and absorbing OM mass. Then the influence of aerosol liquid water on partitioning coefficients could not be compared between these two studies. Thus, in the revised manuscript, we only kept the comparisons of measured gas- and particle-phase concentrations. The partitioning coefficients obtained here were discussed by referring to other studies (e.g., Isaacman-VanWertz et al., 2016; Pye et al., 2018)

Pages 19-20, lines 444-464

“When solubility in aerosol liquid water was considered by assuming a LLPS in ambient aerosols, and whenever WSOM and WIOM partitioned into separate (*Case 2*) or single (*Case 3*) liquid phases, the average $\log K_{p,OM}^m$ of the above mentioned compounds became much closer to or even lay within the range (e.g., levoglucosan) of $\log K_{p,OM}^l$ (Table 1). These results indicated that the aerosol liquid water ($21.3 \pm 24.2 \mu\text{g m}^{-3}$; Table S1) is also an important absorbing phase of ambient polyol tracers in Nanjing. Similarly, the measured average $F\%$ of isoprene SOA tracers in southeastern US and central Amazonia were higher than predictions by assuming instantaneous equilibrium between the gas phase and particulate OM only, and the agreement was improved when parameterization of solubility was included for predictions (Isaacman-VanWertz et al., 2016). But none of these two studies could reasonably predict the temporal variability of $F\%$ or $\log K_{p,OM}^m$. One possible explanation is that the activity coefficients of isoprene SOA tracers and levoglucosan deviate from unity (0.42–2.04; Pye et al., 2018) and vary with PM composition. Pye et al. (2018) re-analyzed the measurement data from Isaacman-VanWertz et al. (2016) using a thermodynamic equilibrium gas-particle partitioning model in two LLPS modes, which involved organic-inorganic interactions and estimations of activity coefficients as a function of liquid PM mixture composition. The resulting predictions captured both the average and diurnal variations of measured $F\%$ for polyol tracers, suggesting a necessity in obtaining time-resolved activity coefficients for the implementation of absorptive equilibrium partitioning model.”

5. Several points in the manuscript, including in the 1st sentence of the abstract, the discussion links gas-particle partitioning to source apportionment. However, gas-particle partitioning has importance in the atmosphere that extends way beyond source apportionment (e.g., it affects the PM mass concentration, the lifetime and distribution of organics in the atmosphere, among others). The study thus has broader relevance than is discussed in the manuscript.

Response:

Gas/particle partitioning is important in predicting the formation, transport, and life time of organic aerosols in the atmosphere, which are mostly involved in atmospheric transport models (e.g., CMAQ).

Tracer-based source apportionment is a different research method highly dependent on the identification of source-specific tracers. As mentioned in the introduction, the target polyols in this work are firstly known as tracers linked with specific emission sources. But their gas-particle partitioning was poorly characterized. Several measurement studies on gas-particle partitioning of polar organic tracers also emphasized its importance in both modeling and tracer-based source apportionment of organic aerosols (e.g., Zhao et al., 2013; Isaacman-VanWertz et al., 2016). That’s the reason gas-particle partitioning of polyol tracers was also linked to source apportionment in some places of the manuscript.

In the abstract, the first sentence has been changed into

“Gas-particle partitioning of water-soluble organic compounds plays a significant role in influencing the formation, transport, and lifetime of organic aerosols in the atmosphere, but is poorly characterized.” (Page 2, lines 31-33)

Technical Corrections

(1) Line 55: “documented” is not the right word here

Response:

Here, “documented for” was replaced by “in”. (Page 4, line 78)

(2) Line 65-66: include “GC-MS” in parentheses after introducing the terms

Response:

It was revised as suggested. (Page 4, line 89)

(3) Line 80-83: I do not follow the logic of this sentence, please clarify

Response:

The original expression has been changed into

“In addition to absorptive partitioning to particulate OM after the formation of oxygenated organic compounds in gas phase, other formation pathways (e.g., reactive uptake) have been identified and are responsible for the large discrepancy between modeled and observed SOA loadings (Jang et al., 2002; Kroll et al., 2005; Perraud et al., 2012).” (Page 5, 104-108)

(4) Line 90: typo in this line

Response:

Thanks, “every” should be “very”. (Page 5, line 117)

(5) Line 93: suggest changing “depict” to “describe”

Response:

It has been changed as suggested. (Page 5, line 120)

(6) Line 101: delete “termed”

Response:

It was deleted as suggested.

(7) Line 116: “unveils” is not the right word here

Response:

It was replaced by “tends to explain”. (Page 6, line 145)

(8) Line 118-120: suggest deleting this sentence

Response:

It was deleted as suggested.

(9) Line 134: change “was” to “were”

Response:

It was changed as suggested. (Page 7, line 163)

(10) Line 153: change “involving” to “using” or similar

Response:

Here, “involving” was replaced by “using”. (Page 7, line 179)

(11) Line 221-222: specify that this is theoretical

Response:

The whole sentence has been deleted in the revised manuscript.

(12) Line 244: edit sentence for grammar

Response:

The original expression has been changed into
“It is therefore suitable to collect gaseous 2-methylterols and levoglucosan using PUF materials only.” (Page 13, lines 312-313)

(13) Line 317: edit sentence for grammar

Response:

The original expression has been changed into
“.....which might be attributed to high levels of vegetation during growing seasons and autumn decomposition (Burshtein et al., 2011).” (Page 17, lines 390-491)

(14) Line 331: delete “data”

Response:

It was deleted and the whole paragraph has been rewritten. (Pages 17-18, lines 397-419)

(15) Line 333: suggest deleting “majorly” and revising the sentence accordingly

Response:

The sentence has been changed into

“Gas-phase C5-alkene triols and 2-methyltetrols had maximum concentrations in summer and significant ($p < 0.05$) day-night variations (Figure S4).....”
(Pages 17, lines 402-403)

(16) Line 360: “prediction” should be plural

Response:

It was changed as suggested. (Page 19, line 451)

(17) Line 380: change “less stable” to “lower”

Response:

We added “lower and” in that sentence. (Page 20, line 478)

(18) Line 386-388: it is quite difficult to follow the discussion here

Response:

Here, the comparisons were reorganized and most of the discussions on $K_{p,OM}$ were rewritten.

“When solubility in aerosol liquid water was considered by assuming a LLPS in ambient aerosols, and whenever WSOM and WIOM partitioned into separate (*Case 2*) or single (*Case 3*) liquid phases, the average $\log K_{p,OM}^m$ of the above mentioned compounds became much closer to or even lay within the range (e.g., levoglucosan) of $\log K_{p,OM}^l$ (Table 1). These results indicated that the aerosol liquid water ($21.3 \pm 24.2 \mu\text{g m}^{-3}$; Table S1) is also an important absorbing phase of ambient polyol tracers in Nanjing.” (Page 19, lines 444-450)

(19) Line 412: efflorescence RH is more relevant than DRH in this scenario

Response:

Here, we deleted the original explanation, and changed the original expression into

“The $\log (K_{H,w}^l/K_{H,e}^m)$ data increased faster when c_{sulfate} approached 0, and deviated from their expected behavior with increased c_{sulfate} . Kampf et al. (2013) selected a threshold c_{sulfate} of 12 mol kg^{-1} ALWC to illustrate the deviation for chamber experiments, and attributed it to elevated viscosity and slow particle-phase reactions at high c_{sulfate} .” (Pages 22, lines 521-525)

(20) Line 462-463: delete “barely” and revise sentence accordingly

Response:

The original expression was changed into

“Then gas-particle partitioning of polyol tracers should have little influence on source apportionment based on particle-phase data in Nanjing.” (Page 24, lines 564-566)

(21) Line 465: not sure what exactly is a “concern”? clarification needed

Response:

The whole sentence was deleted.

(22) Line 478: delete “pre-”

Response:

It was deleted as suggested. (Page 25, line 579)

References

- Ahrens, L., Shoeib, M., Harner, T., Lane, D. A., Guo, R., and Reiner, E. J.: Comparison of annular diffusion denuder and high volume air samplers for measuring per- and polyfluoroalkyl substances in the atmosphere, *Anal. Chem.*, 83, 9622-9628, 10.1021/ac202414w, 2011.
- Isaacman-VanWertz, G., Yee, L. D., Kreisberg, N. M., Wernis, R., Moss, J. A., Hering, S. V., de Sá, S. S., Martin, S. T., Alexander, M. L., Palm, B. B., Hu, W., Campuzano-Jost, P., Day, D. A., Jimenez, J. L., Riva, M., Surratt, J. D., Viegas, J., Manzi, A., Edgerton, E., Baumann, K., Souza, R., Artaxo, P., and Goldstein, A. H.: Ambient gas-particle partitioning of tracers for biogenic oxidation, *Environ. Sci. Technol.*, 50, 9952-9962, 10.1021/acs.est.6b01674, 2016.
- Peters, A. J., Lane, D. A., Gundel, L. A., Northcott, G. L., and Jones, K. C.: A comparison of high volume and diffusion denuder samplers for measuring semivolatile organic compounds in the atmosphere, *Environ. Sci. Technol.*, 34, 5001-5006, 10.1021/es000056t, 2000.
- Pye, H. O. T., Zuend, A., Fry, J. L., Isaacman-VanWertz, G., Capps, S. L., Appel, K. W., Foroutan, H., Xu, L., Ng, N. L., and Goldstein, A. H.: Coupling of organic and inorganic aerosol systems and the effect on gas-particle partitioning in the southeastern US, *Atmos. Chem. Phys.*, 18, 357-370, 10.5194/acp-18-357-2018, 2018.
- Xie, M., Hannigan, M. P., and Barsanti, K. C.: Gas/particle partitioning of n-alkanes, PAHs and oxygenated PAHs in urban Denver, *Atmos. Environ.*, 95, 355-362, <http://dx.doi.org/10.1016/j.atmosenv.2014.06.056>, 2014a.
- Xie, M., Hannigan, M. P., and Barsanti, K. C.: Gas/particle partitioning of 2-methyltetrols and levoglucosan at an urban site in Denver, *Environ. Sci. Technol.*, 48, 2835-2842, 10.1021/es405356n, 2014b.
- Zhao, Y., Kreisberg, N. M., Worton, D. R., Isaacman, G., Weber, R. J., Liu, S., Day, D. A., Russell, L. M., Markovic, M. Z., VandenBoer, T. C., Murphy, J. G., Hering, S. V., and Goldstein, A. H.: Insights into secondary organic aerosol formation mechanisms from measured gas/particle partitioning of specific organic tracer compounds, *Environ. Sci. Technol.*, 47, 3781-3787, 10.1021/es304587x, 2013.
- Zuend, A., and Seinfeld, J. H.: Modeling the gas-particle partitioning of secondary organic aerosol: the importance of liquid-liquid phase separation, *Atmos. Chem. Phys.*, 12, 3857-3882, 10.5194/acp-12-3857-2012, 2012.

This study by Chao Qin et al. reports on filter-based measurements of the gas–particle partitioning of a selection of semi-volatile isoprene oxidation products, levoglucosan and polyols in Nanjing, China. Detailed simultaneous gas and particle phase measurements and assessments of the gas–particle partitioning and influence of aerosol liquid water are relatively scarce. Therefore, this manuscript and the measured data are certainly of interest to the atmospheric chemistry and physics community.

Overall, the manuscript is well written, of adequate lengths and with useful tables and figures. The field sampling and chemical quantification conducted over an extended time span are valuable. The comparison to different predictions by equilibrium gas–particle partitioning models/assumptions is of interest, but it also reveals several issues that need to be addressed.

My main concern is with the provided level of detail on the **uncertainties of the measurements and the theoretical predictions**, as outlined in the general and specific comments below. This manuscript should be (and can be) substantially improved by adding a better discussion and quantification of uncertainties and potential systematic biases as well as clarifications about partitioning mechanisms and involved assumptions. In the present manuscript, the partitioning model discussion is rather confusing, since the title and text suggest a fundamental difference between “absorptive partitioning” and Henry’s law partitioning, not recognizing that Henry’s law is a way of expressing equilibrium (absorptive) gas–liquid partitioning.

Response:

Thanks for the reviewer’s comments, and we will reply these point by point in the reviewers’ specific comments.

General comments

The discussion of the presented mismatch between measured and predicted partitioning of several organic tracers in this manuscript would strongly benefit from a more thorough, quantitative uncertainty analysis of the filter measurements and of the assumptions made with the “theoretical” predictions of partitioning coefficients. This would likely lead to relatively wide error bounds on the median and average partitioning coefficients listed in the tables. At present, the study suggests that there is poor agreement with absorptive (Raoult’s law) partitioning as well as with solubility-based physical Henry’s law partitioning. However, there seem to be substantial uncertainties in the predictions applied and assumptions involved (see the specific comments below).

A comparison to other studies involving the same or similar compounds should be included. The work by Pye et al. (2018) focuses on measurements and conditions in the southeastern United States and includes field measurements and equilibrium partitioning calculations for several polyols and organic acids in common with this study by Qin et al. Pye et al. (2018) also assessed partitioning of 2-methyltetrol, C5 alkene triol, levoglucosan, pinonic acid and other semivolatile compounds. The Pye et al. work includes predicted or assumed liquid–liquid phase separation cases that differ in phase composition from the assumptions made in this study. Importantly, their results show generally a much better agreement between predicted and measured particle phase fractions. Therefore, it is recommended that the authors compare their findings with those from the Pye et al. study and discuss potential reasons for discrepancies in the

partitioning coefficients and their predictability (or that of particle phase fractions).

Response:

In the revised manuscript, we defined three gas-particle partitioning cases. *Case 1* assumed that the particulate OM was the only absorbing material in aerosols based on the equilibrium absorptive partitioning theory. Solubility of polyol tracers in an aqueous phase was included in *Cases 2 and 3*, where water-soluble and water-insoluble OM partitioned into different liquid phases and a single OM phase, respectively. Moreover, measurement uncertainties and their influences on partitioning coefficients of gas versus organic/aqueous phases in aerosols were estimated. It was shown that the average relative uncertainties of measurements and calculated partitioning coefficients ranged from a few percent to ~50%, which corresponded to an uncertainty of less than ± 0.30 for their logarithm values. Although the variability of theoretical partitioning coefficients was large, we still obtained an improved agreement between measurement-based and predicted gas-organic partitioning coefficient ($K_{p,OM}$) for *Cases 2 and 3*. So, aerosol liquid water should have substantial influences on gas-particle partitioning of target polyol tracers in this work.

The interactions of organic and inorganic compounds in the aqueous phase were expected to increase the partitioning of highly water-soluble compounds into the condensed phase (Kroll et al., 2005; Ip et al., 2009; Kampf et al., 2013; Pye et al., 2018). Then the effective Henry's law constants ($K_{H,e}$, mol m⁻³ atm⁻¹) of target polyol compounds in aerosols should be greater than those in pure water ($K_{H,w}$). In the revised manuscript, predicted $K_{H,w}$ values from EPI and SPARC estimates varied by several orders of magnitude. Literature $K_{H,w}$ values were closer to those of 2-methyltetrols and levoglucosan estimated by EPI, while the predicted $K_{H,w}$ with SPARC was unreasonably larger than $K_{H,e}$. So, $K_{H,w}$ values based on EPI estimates were used for further data analysis. Because the "salting-in" effect is a known phenomenon that is not likely linked with a specific physical or chemical mechanism, we made some possible explanations (e.g., reactive uptake) for the enhanced uptake of polyol tracers in aerosol liquid water.

Pye et al. (2018) re-evaluated the measurement data of gas- and particle-phase oxygenated compounds in southeastern US using a thermodynamic equilibrium gas-particle partitioning model in two LLPS modes. The modeling work was based on the AIOMFAC model and programmed inorganic-organic interactions and variations of activity coefficients as a function of liquid PM mixture composition. The resulting predictions captured both the average and diurnal variability of measured $F\%$ for polyol tracers, suggesting a necessity in obtaining time-resolved activity coefficients for the implementation of absorptive equilibrium partitioning model. In this study, partitioning coefficients of polyol tracers were calculated and predicted empirically assuming equilibrium between gas phase and organic/aqueous phases in aerosols. Moreover, particulate OM phase was presumed as an ideal solution in which the activity coefficient ζ_{OM} was equal to 1, and no organic-inorganic interactions were considered. So, the temporal variability of $K_{p,OM}$ was poorly predicted in this work, and the gap between $K_{H,e}$ and $K_{H,w}$ could not be explicitly interpreted. Details of the changes were provided in the revised manuscript and responses to specific comments below.

Specific comments

1. Lines 83 – 87: It is stated that an absorptive partitioning model (which one?) underestimated particle-phase concentrations of carbonyls by several orders of magnitude. Is the argument made by the authors here (from the given phrasing) that absorptive partitioning is an incorrect partitioning mechanism? If so, should it be adsorptive partitioning or what kind? This statement requires further clarification/discussion.

For context, do you mean to say that (1) absorptive partitioning does not take place or (2) that the experiments are not measuring partitioning under equilibrium conditions or that (3) inadequate vapor pressures were used in the partitioning model or (4) something else? For example, could reactive uptake be at play (e.g. mentioned in the cited study by Healy et al., 2008). If the measurement/prediction mismatch is due to reactive uptake, it is questionable to blame absorptive partitioning for this, since that theory may still apply to the parent compound that is partitioning, but further reactions in the particle phase, like hydration of glyoxal, complex formation in presence of sulfate ions or reversible oligomerization may distort the understanding of what species and in what amount is partitioning. It may well be that absorptive equilibrium gas–particle partitioning applies to each of the individual species formed but cannot simply be assumed to be represented by the parent compound considered in the gas phase. Introducing an “effective” Henry’s law coefficient can be used to account for the measured partitioning; however, that formulation is then simply a parameterization and not directly elucidating a physical or chemical mechanism.

Response:

Thanks. Here we did not intend to state that the absorptive partitioning is an incorrect partitioning mechanism. The mismatch between measurement and prediction could be associated with inappropriate assumptions on absorbing phase (e.g., particulate OM only) and formation pathways (e.g., gas-phase oxidation).

To make it clear, we changed the original expression into

“In addition to absorptive partitioning to particulate OM after the formation of oxygenated organic compounds in gas phase, other formation pathways (e.g., reactive uptake) have been identified and are responsible for the large discrepancy between modeled and observed SOA loadings (Jang et al., 2002; Kroll et al., 2005; Perraud et al., 2012). Unlike non-polar species (e.g., *n*-alkanes, polycyclic aromatic hydrocarbons) and alkanolic acids that are well simulated (Simcik et al., 1998; Xie et al., 2014a; Yatavelli et al., 2014; Isaacman-VanWertz et al., 2016), particle-phase concentrations of carbonyls were underestimated by several orders of magnitude when particulate OM is presumed as the only absorbing phase in ideal condition (Healy et al., 2008; Kampf et al., 2013; Shen et al., 2018).” (Page 5, lines 104-114)

In the original manuscript, the absorptive partitioning means partitioning between gas phase and particulate OM in aerosols; Henry’s law partitioning corresponds to the equilibrium between gas phase and aerosol liquid water. We have clarified this throughout the manuscript.

2. Line 89: “favored the formation of pinonaldehyde”; do you mean “partitioning” instead of “formation”? The formation of pinonaldehyde (in the gas phase) is likely independent from aerosol water content.

Response:

Zhao et al. (2013) stated that the aerosol water plays a role in the formation of particle-phase pinonaldehyde in the atmosphere. This might be related to the water uptake. So, here we replaced “favored” with “played a role in”. (Page 5, lines 116)

3. Line 100: The work by Volkamer et al. (2009) on effective Henry’s law partitioning and aqueous phase chemistry could also be cited here and perhaps discussed in context of the findings from this study later in the article.

Response:

The work by Volkamer et al. (2009) has been cited here.

“An effective Henry’s law coefficient ($K_{H,e}$, mol m⁻³ atm⁻¹) can be used to account for the measured partitioning between the gas phase and aerosol liquid water (Volkamer et al., 2009).” (Page 5-6, lines 124-126).

Volkamer et al. (2009) focused on the effect of seed chemical composition and photochemistry on SOA yields. They found that the WSOC photochemical reactions can cause increased SOA yield. So, this work was also cited later in the manuscript.

“Moreover, log ($K_{H,w}^t/K_{H,e}^m$) values of polyol tracers also negatively correlated with the aqueous-phase concentrations of WSOC (c_{WSOC} , Figure S6), but not NH₄⁺ or NO₃⁻. This dependence might be associated with the “like-dissolves-like” rule, or indicate the importance of aqueous-phase heterogeneous reactions (Hennigan et al., 2009; Volkamer et al., 2009).” (Pages 23-24, lines 552-556)

4. Line 133 – 135: From this description of the gas and aerosol measurements using filters in series, it is not clear how much the uptake of gaseous (semivolatile) organic compounds on accumulated aerosol mass loading of filter 1 (Qf) will contribute to the total concentration on the particle filter. Based on absorptive equilibrium partitioning theory, the accumulated condensed-phase aerosol mass on the first filter may shift the actual gas–particle partitioning in the ambient air to favor additional partitioning from the gas phase to the condensed phase on the filter while the sampling flow passes through the filter, thus possibly leading to a systematic particle phase mass concentration bias. Given the long sampling times, this may constitute a substantial bias. Were such potential issues quantified in controlled experiments? Please discuss.

Response:

Based on absorptive equilibrium partitioning theory, the partitioning coefficient of a certain compound between gas and particulate OM phases ($K_{p,OM}$, m³ ug⁻¹) is defined as

$$K_{p,OM} = \frac{F/M_{OM}}{A} \quad (1)$$

where F and A are (ng m⁻³) are particle- and gas-phase concentrations, and M_{OM} (μg m⁻³) is the mass concentration of particulate OM. It is assumed that the quantity F/M_{OM} (ng μg⁻¹) represents the equilibrium concentration in the particulate matter (Pankow and Bidleman, 1992; Liang et al., 1997).

In previous studies, $K_{p,OM}$ values based on offline measurements were typically obtained using sampling periods of many hours (e.g., 8, 12, or 24 h). When ambient concentrations (F , A , or M_{OM}) or temperature change within a sampling interval, the particulate OM initially collected on the filter will tend to re-equilibrate with the A value through evaporation or absorption. Then whether the accumulated aerosol mass will uptake or release gaseous organic compounds depends on how changes in F , A , M_{OM} , and ambient temperature take place. Measured values of F , A , and M_{OM} will be averages over the whole sampling period, not reflecting real-time atmospheric concentrations. Therefore, the situation raised in the comment seems not applicable to this study.

5. Line 152 – 154: “Concentrations of aerosol liquid water were predicted by ISORROPIA II model”; this prediction will only account for water uptake by inorganic ions but neglect any water uptake by hygroscopic organic compounds (such as some WSOC), right? It may therefore lead to an underestimation of the WSOC effect on organic partitioning. The authors could use a simple estimation based on typical organic hygroscopicity parameters (κ) and the median or actual RH values to estimate the organic-contributed water content by the WSOC mass fraction in particles.

Response:

According to Isaacman-VanWertz et al. (2016), the water uptake by WSOC (W_O , $\mu\text{g m}^{-3}$) could be estimated as

$$W_O = \frac{V_{\text{WSOC}} \times [\kappa \times (\text{O:C})]}{\left(\frac{100}{\text{RH}\%} - 1\right)} \quad (2)$$

where V_{WSOC} represents WSOC volume, and is calculated as the organic mass (WSOC $\times 1.6$) divided by its density (1.4 g cm^{-3}). The hygroscopicity parameter (κ) and oxygen to carbon ratio (O:C) of WSOC were assumed as 0.10 and 0.5, respectively, based on field and laboratory studies (Taylor et al., 2017; Cai et al., 2020). The resulting W_O had an average of $0.47 \pm 1.14 \mu\text{g m}^{-3}$, far below the amount caused by inorganic ions ($21.3 \pm 24.2 \mu\text{g m}^{-3}$). Taylor et al. (2017) predicted a growth factor range of 1.00–1.20 with κ varying from 60 to ~100%, which lead to a comparable average W_O ($0.42 \pm 0.70 \mu\text{g m}^{-3}$) in this work. Thus, the water content contributed by WSOC was not accounted for in this work.

We have clarified this in the revised manuscript and supplementary information. (Page 8, lines 180-182)

“The estimated water content contributed by hygroscopic WSOC was relatively small ($< 1 \mu\text{g m}^{-3}$) and not accounted for in this work (Text S1 of supplementary information).”

6. Line 209: I suggest adding these equations to the main text.

Response:

We have defined three partitioning cases and included these equations in the revised manuscript. (Pages 10 -13, lines 238-295)

“Here, we defined three partitioning cases to explore the influence of dissolution in aerosol liquid water on gas-particle partitioning of polyol tracers in the atmosphere. *Case 1* presumes instantaneous equilibrium between the gas phase and particulate OM

based on the equilibrium absorptive partitioning theory. In this case, particulate OM is assumed to be the only absorbing phase and behave as an ideal solution. Then the absorptive gas-particle partitioning coefficients ($K_{p,OM}$, $m^3 \mu g^{-1}$) were calculated from measurements ($K_{p,OM}^m$) and predicted theoretically ($K_{p,OM}^t$) as follows

$$K_{p,OM}^m = \frac{F/M_{OM}}{A} \quad (2)$$

$$K_{p,OM}^t = \frac{RT}{10^6 \overline{MW}_{OM} \zeta_{OM} p_L^o} \quad (3)$$

where M_{OM} denotes the mass concentration of absorptive organic matter ($OM = OC \times 1.6$; Turpin and Lim, 2001); F ($ng\ m^{-3}$) and A ($ng\ m^{-3}$) are particulate and gaseous concentrations of individual polyols, respectively. In eq 2, R ($m^3\ atm\ K^{-1}\ mol^{-1}$) and T (K) are the ideal gas constant and ambient temperature; \overline{MW}_{OM} , average molecular weight of absorptive OM, is set at $200\ g\ mol^{-1}$ for all samples (Barsanti and Pankow, 2004; Williams et al., 2010); ζ_{OM} denotes the mole fraction scale activity coefficient, and is presumed to be unity for all species in each sample; p_L^o (atm) is the vapor pressure of each pure compound, and is predicted with several estimation tools and adjusted for each sampling interval based on the average temperature (Text S3 and Table S4).

Due to the influence of mixing state and water content in aerosols, several studies modeled the gas-particle partitioning of oxygenated organic compounds by defining a liquid-liquid phase separation (LLPS) in the aerosol (Zuend and Seinfeld, 2012; Pye et al., 2018). The organic-inorganic interactions and changes of activity coefficients in aqueous mixtures were fully considered as well. In this study, we proposed a simplified LLPS partitioning mechanism (*Case 2*) in Figure 1. First, aerosol water and water-insoluble OM ($WIOM = OM - WSOC \times 1.6$) exist in two separate liquid phases, and WSOC and inorganic ions are totally dissolved in the aqueous phase. The distribution of polyol tracers between aqueous (F_W , $ng\ m^{-3}$) and WIOM (F_{WIOM} , $ng\ m^{-3}$) phases is simply depicted by their octanol-water partition coefficients (K_{OW})

$$K_{OW} = \frac{F_{WIOM}/V_{WIOM}}{F_W/V_W} = \frac{c_{WIOM}}{c_W} \quad (4)$$

where V_{WIOM} and V_W are volumes (m^3) of WIOM and water in aerosols per cubic meter air; c_{WIOM} and c_W are solution concentrations ($ng\ m^{-3}$) of polyols concentrations in organic and aqueous phases; $\log K_{OW}$ values of target polyols were estimated using the Estimation Programs Interface (EPI) Suite developed by the US Environmental Protection Agency and Syracuse Research Corporation (Table S4; US EPA, 2012). The density of organic matter and water (ρ_w) in aerosols are set at 1.4 and $1.0\ g\ cm^{-3}$, respectively (Isaacman-VanWertz et al., 2016; Taylor et al., 2017). Second, gas-phase polyol tracers are in equilibrium with hydrophobic OM and the aqueous phase, respectively

$$K_{p,WIOM}^m = \frac{F_{WIOM}/M_{WIOM}}{A} \quad (5)$$

$$K_{H,e}^m = \frac{\frac{F_W}{M_i}}{\frac{A}{M_i} \times R \times T \times \frac{c_{ALW}}{\rho_w}} = \frac{\rho_w \times F_W}{A \times R \times T \times c_{ALW}} \quad (6)$$

where $K_{H,e}^m$ ($mol\ m^{-3}\ atm^{-1}$) is the measurement-based effective Henry's law coefficient; M_{WIOM} represents the mass concentration ($\mu g\ m^{-3}$) of WIOM; M_i ($g\ mol^{-1}$) is the molecular weight of specific compound; c_{ALW} ($\mu g\ m^{-3}$) is the mass concentration of aerosol liquid water predicted using ISORROPIA II model. *Case 3* is generally the same as *Case 2*, and the only difference is that water-soluble OM (WSOM) and WIOM exist in a single organic phase. Here total particulate OM was used instead of WIOM to assess the distribution of polyol tracers between aqueous and organic phases, and

calculate partitioning coefficients of gas vs. particulate organic ($K_{p,OM}^m$) and aqueous ($K_{H,e}^m$) phases. Note that the polarity of particulate OM phase in *Case 3* was expected to increase, then using K_{OW} to calculate the distribution of polyols between organic and aqueous phases might lead to underestimated $K_{p,OM}^m$ and overestimated $K_{H,e}^m$. For comparison purposes, the Henry's law coefficient in pure water at 25 °C ($K_{H,w}^*$) was estimated using EPI and SPARC (Hilal et al., 2008; <http://archemcalc.com/sparc-web/calc>), respectively (Table S4), and was adjusted for each sampling interval due to the changes in ambient temperature using van 't Hoff equation (Text S4)."

7. Line 212: If I understand your procedure, the ISORROPIA-derived aerosol water content is not accounting for water associated with WSOC, which could be substantial at high RH and when the WSOC represent a significant mass fraction of aerosol. Also, actual interactions among organics and ions within particle phases may affect the partitioning (both between liquid phases and gas/particle), which I assume is ignored in this work. Furthermore, WSOC, while water-extractable by definition, can be of relatively moderate polarity and may preferably partition to the WIOM organic-rich phase in presence of dissolved salts in an aqueous phase (see e.g. Zuend et al., 2012; You et al., 2014; Pye et al. 2018). Hence, it would be useful to estimate errors from such effects on the determined KOW. It may also be adequate to consider other liquid-liquid phase separation scenarios, such as assuming that all WIOM and WSOC organics partitioned to one aqueous organic phase and all inorganic salts to a separate aqueous inorganic phase (compare to Fig. 3 of Pye et al., 2018).

Response:

As mentioned in responses to *Comment 5*, the estimated contribution of WSOC to aerosol liquid water is relatively small. In this study, gas-particle partitioning coefficients of polyol tracers were calculated and predicted empirically by assuming equilibrium between gas phase and organic/aqueous phases in aerosols, and the organic-inorganic interactions were not considered. This might be an important reason for the gap between measurement-based $K_{H,e}$ and predicted $K_{H,w}$. Pye et al. (2018) re-analyzed the measurement data from Isaacman-VanWertz et al. (2016) using a thermodynamic equilibrium gas-particle partitioning model in two LLPS modes, which involved inorganic-organic interactions and estimations of activity coefficients as a function of liquid PM mixture composition. The resulting predictions captured both the average and diurnal variations of measured $F\%$ for polyol tracers, suggesting a necessity in obtaining time-resolved activity coefficients for the implementation of absorptive equilibrium partitioning model. These discussions have been added in the revised manuscript. (Pages 19-20, lines 457-464)

As suggested by the reviewer, we defined a third partitioning case assuming all WIOM and WSOC organics partitioned to a single organic phase (*Case 3*, see responses to *Comment 6*). However, the measurement-based partitioning coefficients of polyol tracers in *Case 3* are very close to those in *Case 2*, where WIOM and WSOC were assumed to partition into separate liquid phases. In the revised manuscript, we defined three gas-particle partitioning cases and re-analyzed the difference between measured and predicted partitioning coefficients. (Pages 10-13, lines 238-295, see responses to *Comment 6*; Sections 3.4 and 3.5 in the revised manuscript, Tables 1 and 2 below).

Table 1. Comparisons of measurement-based $\log K_{p,OM}$ ($\text{m}^3 \mu\text{g}^{-1}$) at three proposed cases and predicted values.

Species	No. of obs.	Log $K_{p,OM}^a$			Log $K_{p,OM}^b$			
		Case 1	Case 2	Case 3	EPI	EVAPORATION	SPARC	SIMPOL
Isoprene SOA tracers								
C5-alkene triol 1	53	0.33 ± 0.71	-0.79 ± 0.86	-0.82 ± 0.85	-3.09	-2.84	-1.19	-2.88
C5-alkene triol 2	63	0.15 ± 0.55	-1.02 ± 0.74	-1.05 ± 0.73	-3.62	-3.67	-4.14	-2.85
C5-alkene triol 3	83	0.35 ± 0.68	-0.83 ± 0.86	-0.86 ± 0.85	-2.90	-2.65	-1.00	-2.69
2-Methylthreitol	101	-0.12 ± 0.48	-2.09 ± 0.71	-2.09 ± 0.70	-1.87	-1.30	-1.18	-0.47
2-Methylerythritol	95	-0.011 ± 0.58	-1.96 ± 0.71	-1.96 ± 0.71	-1.90	-1.34	-1.22	-0.50
Biomass burning tracer								
Levogluconan	65	2.23 ± 0.72	0.63 ± 0.90	0.62 ± 0.90	-0.04	-0.81	1.04	-0.76
Sugars and sugar alcohols								
Meso-erythritol	31	0.87 ± 0.53	-1.43 ± 0.60	-1.43 ± 0.60	-0.65	-1.21	-0.45	
Fructose	85	0.65 ± 0.73	-1.20 ± 0.83	-1.20 ± 0.89	1.17	2.76	6.94	
Mannose	74	0.62 ± 0.71	-2.12 ± 0.95	-2.12 ± 0.95	1.28	2.13	4.77	
Glucose	88	0.42 ± 0.67	-2.77 ± 0.93	-2.77 ± 0.93	0.34	3.75	7.32	
Xylitol	22	0.24 ± 0.54	-2.61 ± 0.72	-2.61 ± 0.72	3.37	2.34	3.57	
Arabitol	30	1.46 ± 0.89	-1.35 ± 1.24	-1.35 ± 1.24	3.25	1.67	2.90	
Manitol	65	1.08 ± 0.63	-2.24 ± 0.95	-2.24 ± 0.95	2.33	4.16	6.68	

^a Average ± standard deviation; ^b temperature range: -4–36 °C.

Table 2. Comparisons of measurement-based $\log K_{H,e}$ ($\text{mol m}^{-3} \text{atm}^{-1}$) and predicted $\log K_{H,w}$ of individual polyol tracers.

Species	No. of obs.	Log $K_{H,e}$ (Cases 2) ^a			Log $K_{H,w}$ ^b	
		Median	Average	Range	EPI	SPARC
Isoprene SOA tracers						
C5-alkene triol 1	53	14.0	13.9 ± 0.86	11.5 – 16.4	7.22	11.7
C5-alkene triol 2	63	13.7	13.6 ± 0.73	11.2 – 16.1	7.34	7.66
C5-alkene triol 3	83	13.9	13.8 ± 0.85	10.6 – 16.1	7.43	11.9
2-Methylthreitol	101	13.4	13.3 ± 0.70	10.9 – 14.8	10.0	14.1
2-Methylerythritol	95	13.5	13.5 ± 0.71	11.6 – 15.6	9.95	14.1
Biomass burning tracer						
Levogluconan	65	15.7	15.7 ± 0.90	13.2 – 17.3	13.4	16.1
Sugars and sugar alcohols						
Meso-erythritol	31	14.5	14.4 ± 0.60	12.8 – 15.6	9.65	13.8
Fructose	85	14.2	14.1 ± 0.89	11.9 – 16.5	14.7	19.9
Mannose	74	14.0	14.1 ± 0.94	12.1 – 16.8	10.9	18.8
Glucose	88	13.9	13.9 ± 0.93	11.3 – 16.3	14.7	20.9
Xylitol	22	13.8	13.7 ± 0.72	12.6 – 15.0	12.1	18.1
Arabitol	30	15.1	15.0 ± 1.23	13.0 – 18.2	11.3	17.4
Mannitol	65	14.6	14.5 ± 0.94	12.1 – 16.4	12.9	20.8

^a Log $K_{H,e}$ values of Case 3 had ignorable difference, and were not exhibited separately; ^b temperature range: -4–36 °C.

8. Lines 221 – 222: The rather low octanol–water partitioning coefficients indicate not only better solubility in water but also that the polyols of moderate to high polarity have low solubility in octanol; this is because octanol is a rather low polarity medium as choice for representing organic aerosol. SOA-rich phases may be of substantially higher polarity than octanol yet still form a separate phase from an aqueous salt-rich phase (e.g. You et al., 2014). This should be acknowledged, and consequences of partitioning assumptions considered in the uncertainty analysis.

Response:

As we mentioned in responses to the general comment and specific comment 6, three partitioning cases were proposed in the revised manuscript. In the newly defined *Case 3* where WIOM and WSOC organics partitioned to a single organic phase, the solubility of polyol tracers in the organic phase was expected to increase, and using K_{OW} to calculate the distribution between organic and aqueous phases in aerosols would lead to an underestimation of $K_{p,OM}$, which might not be reasonably adjusted. This was acknowledged in the revised manuscript.

Pages 12-13, lines 288-291.

“Note that the polarity of particulate OM phase in *Case 3* was expected to increase, then using K_{OW} to calculate the distribution of polyols between organic and aqueous phases might lead to underestimated $K_{p,OM}^m$ and overestimated $K_{H,e}^m$.”

9. Line 224 and SI Eq. (2), Text S2: In the SI, it is stated that for the absorptive partitioning prediction an average organic molar mass MW_{OM} of 200 g/mol was used. This seems to be a common and reasonable assumption, but only for a water-free organic absorbing phase. However, for the partitioning of WSOC compounds when assumed to prefer the aqueous phase, one should account for the low molar mass of water present in substantial amounts in that phase, which would lower the weighted mean molar mass significantly (Liu et al., 2021; Gorkowski et al. 2019). Please consider this and, where applicable, correct the estimated partitioning coefficients.

Response:

Among the three proposed partitioning cases in the revised manuscript, *Case 1* assumed that the particulate OM was the only absorbing material in aerosols, ignoring the influence of aerosol liquid water. While in *Cases 2* and *3*, particulate OM and aerosol liquid water were assumed to exist in separate liquid phases, and there was no phase distribution of water. By comparing measurement-based and predicted partitioning coefficients for different cases, we inferred that the aerosol liquid water should play a significant role in influencing gas-particle partitioning of polyol tracers (Table 1, see responses to *Comment 7*). An enhanced uptake of polyol tracers was also identified, which should be closely associated with the organic-inorganic interactions.

In this study, the particulate organic phase was assumed to contain no water, and partitioning of gaseous polyols to organic and aqueous phases in aerosols were assessed separately. So, the influence of water content on molecular weight of particulate organic matter (MW_{OM}) was not considered.

10. Line 233: Use of the EPI suite estimations should be considered uncertain by about one order of magnitude (or more in certain cases) for predictions involving multifunctional semivolatile compounds. A comparison to other estimation methods for physical Henry's law constants (and their estimated uncertainties) may provide some information on the reliability of this method.

Response:

Thanks. We also obtained Henry's law constants of polyols in pure water at 25 °C

($K^*_{H,w}$) from SPARC (Hilal et al., 2008; <http://archemcalc.com/sparc-web/calc>) estimates and literatures (Table S4).

Although the $K^*_{H,w}$ ($\text{mol m}^{-3} \text{atm}^{-1}$) from EPI and SPARC differed by several orders of magnitude, literature values of isoprene SOA and levoglucosan were closer to the estimates of EPI (Table S4). If SPARC $K^*_{H,w}$ values were used, the average $\log K^m_{H,e}$ of most polyol tracers would be lower than predictions ($\log K^l_{H,w}$, Table 2; see responses to *Comment 7*), indicating that the aqueous phase of ambient aerosol is less hospitable to polyol tracers than pure water. This is in conflict with the fact that the interactions of organic compounds, water, and inorganic ions in aerosols will increase the partitioning of highly oxygenated compounds ($\text{O:C} \geq 0.6$; e.g., isoprene SOA tracers and levoglucosan) into the particle phase (Pye et al., 2018). Several studies identified a close relationship between salt concentrations of aerosol water and enhanced uptake of very polar compounds (Kroll et al., 2005; Ip et al., 2009; Kampf et al., 2013). Thus, $\log K^l_{H,w}$ values of EPI estimates were used for further data analysis. (Pages 20-21, lines 482-494)

11. Line 328 – 332: Why was a linear regression/relationship used? Partitioning theory would suggest that it should be a sigmoidal relationship (if applicable), e.g. O’Meara et al (2014); Donahue et al. (2009). However, the partitioning of a specific compound will also depend on the condensed phase absorbing mass concentration (in organic or aqueous phase, as appropriate) and on non-ideality, such as presence of phase separation. Given that only particle phase fraction data above ~80% were determined from the measurements, the expected sigmoidal relationship is perhaps not clear from the data alone.

Response:

Thanks. The original Figure 2 (now Figure 3) and associated discussions was changed considering the uncertainties in measurements

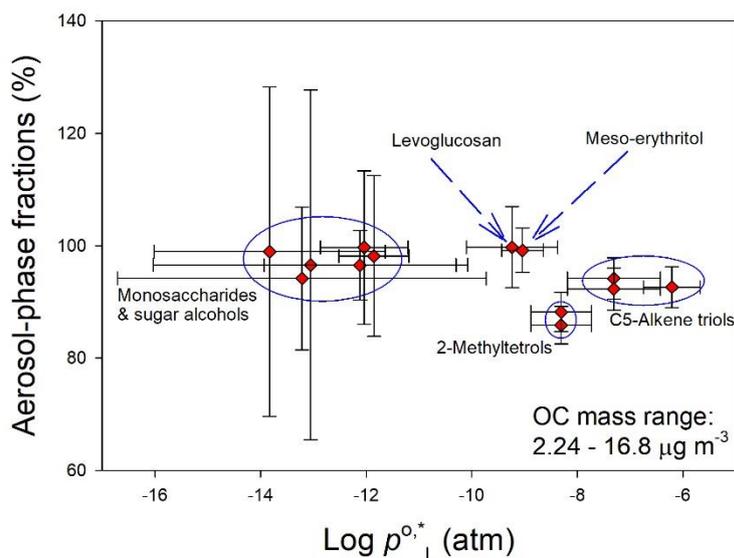


Figure 3. Average particle-phase fractions and $\log p^{0,*}_L$ of individual polyol tracers. Whiskers represent uncertainties of $F\%$ and one standard deviation of $\log p^{0,*}_L$ derived from different estimation tools.

Pages 17-18, lines 410-419

“In Figure 3, the average $F\%$ uncertainties (6.16–31.2%) of monosaccharides (e.g., fructose) and sugar alcohols (e.g., mannitol) were larger than those of isoprene SOA tracers and levoglucosan (3.33–7.24%) due to their low and variable recoveries (Table S2) and excessive breakthrough (Figure S2). However, the estimated uncertainties of $F\%$ for less volatile polyols ($p^{0,*}_L < \sim 10^{-10}$ atm) were not physically meaningful, as more than 95% of these compounds existed in the particle phase. Considering the uncertainties in $F\%$ and $\log p^{0,*}_L$ and high average $F\%$ ($> 85\%$) of target polyol tracers, a dependence of $F\%$ on the vapor pressure could not be determined, and the seasonality and day-night difference ($p > 0.05$) of $F\%$ were obscured.”

12. Line 334: “their $F\%$ values did not show seasonality or day-night difference”;

The range of particle phase fractions observed may not allow for such conclusions if the material is predominantly in the particle phase. Uncertainties in the measurements and temperature dependence of the vapor pressures may mask actual variations.

Response:

The original expression has been changed. See responses to *Comment 11* (Pages 17-18, lines 410-419).

13. Lines 339 – 340: “Thus, the changes in vapor pressures with the ambient temperature might not be the main factor driving gas-particle partitioning of polyol tracers in northern Nanjing.”

What about variations in organic aerosol mass concentrations as additional influence?

Response:

Thanks. Xie et al. (2014) found that the gas-particle partitioning of 2-methyltetrols and levoglucosan in urban Denver were dependent on the variations in ambient temperature and absorbing organic matter (MOM). So, the original expression was changed into

“Thus, the changes in vapor pressures with the ambient temperature and/or particulate OM loadings might not be the main factors driving gas-particle partitioning of polyol tracers in Nanjing.” (Page 17, lines 408-410)

14. Line 360: The re-evaluation of the SV-TAG measurements by Isaacman-VanWertz et al. (2016) in the study by Pye et al. (2018) (see their Fig. 5) involving other models, considerations of vapor pressure adjustments and additional measurement comparisons, shows that higher and lower particle phase fractions were predicted, but that generally the agreement between models and observed $F\%$ were consistent across a selection of tracers and much better than the orders of magnitude differences reported in this manuscript.

Response:

When the dissolution of polyols in aerosol water was included, the agreement

between measurement-based and predicted partitioning coefficients of gas vs organic phases ($K_{p,OM}$) were improved substantially. However, the variability of gas-particle partitioning was poorly predicted. The study by Pye et al. (2018) was cited and discussed to explain the discrepancy between measurements and predictions. (Pages 19-20, lines 450-464)

“Similarly, the measured average $F\%$ of isoprene SOA tracers in southeastern US and central Amazonia were higher than predictions by assuming instantaneous equilibrium between the gas phase and particulate OM only, and the agreement was improved when parameterization of solubility was included for predictions (Isaacman-VanWertz et al., 2016). But none of these two studies could reasonably predict the temporal variability of $F\%$ or $\log K_{p,OM}^m$. One possible explanation is that the activity coefficients of isoprene SOA tracers and levoglucosan deviate from unity (0.42–2.04; Pye et al., 2018) and vary with PM composition. Pye et al. (2018) re-analyzed the measurement data from Isaacman-VanWertz et al. (2016) using a thermodynamic equilibrium gas-particle partitioning model in two LLPS modes, which involved organic-inorganic interactions and estimations of activity coefficients as a function of liquid PM mixture composition. The resulting predictions captured both the average and diurnal variations of measured $F\%$ for polyol tracers, suggesting a necessity in obtaining time-resolved activity coefficients for the implementation of absorptive equilibrium partitioning model.”

15. Lines 364 – 367: These statements are misleading and need to be rephrased. Henry's law partitioning is a form of absorptive partitioning (in contrast to adsorptive partitioning). In the case of SVOCs and LVOCs, the difference between vapor–liquid equilibrium and liquid-phase mixing described by using Raoult's law or Henry's law (when accounting for non-ideal mixing) is essentially a matter of choice of reference state (while for non-vapor gases only Henry's law can be applied).

The observed large differences between measurements and different predictions could be the result of a combination of issues and uncertainties associated with the measurements and the models used. If reactive uptake is considered to be the key difference between predictions and measurements, this should be clarified.

Response:

Thanks. In the revised manuscript, the absorptive partitioning of gaseous polyols to organic and aqueous phases in aerosols were clearly distinguished, and uncertainties of measurements and predictions and their influences on the calculation of partitioning coefficients were estimated.

Through the comparisons of measurement-based and predicted $K_{p,OM}$ before and after the inclusion of polyols dissolution in aerosol water, we inferred that the aerosol liquid water is also an important absorbing phase of ambient polyol tracers in Nanjing. The large gaps of $K_{H,e}^m$ versus $K_{H,w}^t$ could be partly parameterized using the equation defining “salting-in” effects. However, the “salting-in” effect is a known phenomenon that is not likely linked with a specific physical or chemical mechanism. According to existing studies, reactive uptake, aqueous phase reactions, and chemical similarity between partitioning species and the absorbing phase might be responsible for increasing the partitioning of polyol tracers into the condensed phase. In the revised manuscript, we have re-organized and re-written most of the discussions on comparisons between measured and predicted partitioning coefficients. (Sections 3.4

and 3.5, Pages 18-23, lines 420-560).

Here are some discussions on mechanisms related to the gap between $K_{H,e}^m$ and $K_{H,w}^t$. (Pages 23-24, lines 534-560)

“However, the “salting-in” effect is a known phenomenon that is not likely linked with a specific physical or chemical mechanism. Quantum chemical calculation results indicated negative Gibbs free energy of water displacement for interactions between SO_4^{2-} and glyoxal monohydrate (Waxman et al., 2015). The net “salting-in” effect of 1-nitro-2-naphthol in NaF solution was interpreted by postulating hydrogen bonding (Almeida et al., 1983). A direct binding of cations to ether oxygens was proposed to be responsible for the increased solubility of water-soluble polymers (Sadeghi and Jahani, 2012). Due to the moderate correlations and negative intercepts in Figures 4 and S5, the gap between $K_{H,e}^t$ and $K_{H,w}^m$ cannot be closed by the “salting-in” effect alone. Shen et al. (2018) also obtained negative intercepts when plotting $\log(K_{H,w}^t/K_{H,e}^m)$ over c_{sulfate} for glyoxal and methylglyoxal in ambient atmosphere, and attributed this to unknown gas-particle partitioning mechanisms. Evidences showing that the thermal degradation of less volatile oligomers and organosulfates can lead to an overestimation of 2-methyltetrols by 60–188% when using a conventional GC/EI-MS method (Cui et al., 2018). To fit the gas-particle distribution of 2-methyltetrols in southeastern US, 50% of particulate 2-methyltetrols was presumed to exist in chemical forms with much lower vapor pressures by Pye et al. (2018). So, the reactive uptake and aqueous phase chemistry could be explanations for the enhanced uptake of isoprene SOA tracers. Moreover, $\log(K_{H,w}^t/K_{H,e}^m)$ values of polyol tracers also negatively correlated with the aqueous-phase concentrations of WSOC (c_{WSOC} , Figure S6), but not NH_4^+ or NO_3^- . This dependence might be associated with the “like-dissolves-like” rule, or indicate the importance of aqueous-phase heterogeneous reactions (Hennigan et al., 2009; Volkamer et al., 2009). Although several studies have estimated Henry’s law constants for a variety of polar organic compounds in pure water (e.g., polyols and polyacids; Compornolle and Müller, 2014a, b), more work is warranted to decrease the estimation uncertainty and explain their increased partitioning toward aerosol liquid water explicitly.”

16. Lines 387 – 389: The statements on these lines seem to support the conclusion that absorptive partitioning may be applicable to describing the partitioning of these isoprene SOA tracers, but only if one uses the “appropriate” absorbing organic phase mass in the estimation of the measured K_p values (and given the uncertainty in the vapor pressures and activity coefficients, this seems to be reasonable). The phrasing could be improved to make that point.

Response:

Yes, the agreement between measurement-based and predicted $K_{p,OM}$ will be significantly improved when the “appropriate” absorbing organic phase is used.

To make it clear, we defined three partitioning cases in the revised manuscript. *Case 1* assumed that the particulate OM was the only absorbing material in aerosols based on the equilibrium absorptive partitioning theory. Solubility of polyol tracers in an aqueous phase was included in *Cases 2 and 3*, where water-soluble and water-insoluble OM partitioned into different liquid phases and a single OM phase, respectively. We found a much better agreement between measurement-based and predicted $K_{p,OM}$ after considering polyols dissolution in aerosol liquid water, indicating

that aerosol liquid water is also an important absorbing phase of ambient polyols in Nanjing.

Pages 18-19, lines 421-450

“To understand if particulate OM is the only absorbing phase in aerosols for polyol tracers in Nanjing, the absorptive partitioning coefficients of gas vs. organic phases were calculated based on measurement results ($K_{p,OM}^m$) for predefined *Cases 1-3* and predicted theoretically ($K_{p,OM}^t$) using eq. 3 and vapor pressures listed in Table S4. In Table 1, $K_{p,OM}^t$ ranges of isoprene SOA tracers, levoglucosan, and meso-erythritol are within two orders of magnitude, while those of monosaccharides and mannitol are larger ($> 10^3$). When particulate OM was assumed as the only absorbing phase (*Case 1*), the average $K_{p,OM}^m$ of isoprene SOA tracers, levoglucosan, and meso-erythritol were more than 10 times greater than most of their $K_{p,OM}^t$ (Table 1), and this difference was not likely susceptible to measurement uncertainties. As shown in Table S5, the average relative uncertainties of measurement-based partitioning coefficients are all $< 50\%$, leading to an uncertainty of $\log K_{p,OM}^m$ less than ± 0.30 ..

When solubility in aerosol liquid water was considered by defining a LLPS in ambient aerosols, and whenever WSOM and WIOM partitioned into separate (*Case 2*) or single (*Case 3*) phases, the average $\log K_{p,OM}^m$ of the above mentioned compounds became much closer to or even lay within the range (levoglucosan) of $\log K_{p,OM}^t$ (Table 1). These results indicated that the aerosol liquid water ($21.3 \pm 24.2 \mu\text{g m}^{-3}$; Table S1) is also an important absorbing phase of ambient polyol tracers in Nanjing.”

17. Lines 415 – 421: The finding that the intercept in Fig. 3 of the linear regression does not go through 0.0 indicates that there are substantial uncertainties, making this comparison far less convincing. The scatter in the data is large, also hinting at salting-in as an effect alone does not seem to be a good explanation of the deviations between predicted and measurement-derived Henry's law partitioning. The authors also mention this on lines 445 – 448. There may be other confounding factors that happen to correlate with sulfate concentration; leading to a spurious conclusion of a causal salting-in effect that is not strongly supported by the provided evidence. For example, the ratio of WIOC to WSOC organic material may correlate with sulfate concentrations since sulfate and ammonia amounts will affect and respond to aerosol pH, which may also correlate with RH and absolute ALWC (Pye et al., 2020). Did the authors consider this?

Furthermore, a salting-in of polyols by sulfate is a finding that would be contrary to other studies on liquid–liquid phase partitioning involving polyols and ammonium sulfate, e.g. see Table 1 of Marcolli and Krieger (2006). In the study by Marcolli and Krieger (2006), ammonium sulfate led to salting-out while ammonium nitrate was found to show a salting-in effect on polyols. However, the complexity of the samples from Nanjing, where perhaps acidity and other aerosol components affect uptake, may differ from those in laboratory experiments by Marcolli and Krieger. Please discuss your findings of potential reasons for the model–measurement discrepancies and sulfate influence also in comparison to findings on salting-in/out from those studies.

Response:

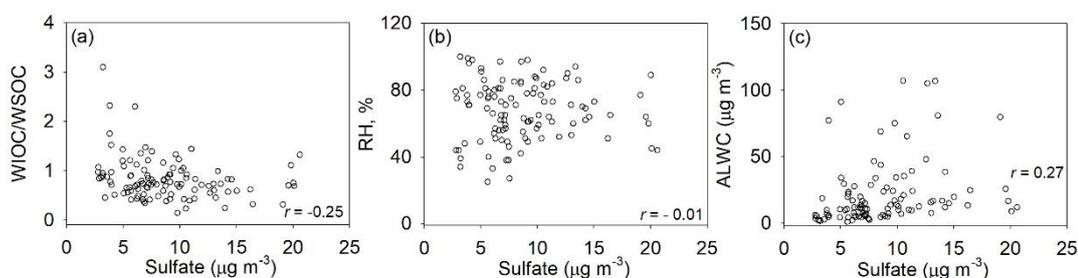
Although the “salting-in” effect has been known for a long time, it is poorly characterized at high salt concentrations and is not understood mechanistically. The

“salting-in” effect could be considered as a phenomenon that is not specifically linked with a physical or chemical mechanism. So, the equation defining “salting-in” effects was only used to parameterize the enhanced uptake of polyols in aerosol liquid water. In the revised manuscript, we provided some guesses on mechanisms related to the “salting-in” effect from previous work and other explanations for increased partitioning to the particle phase.

Pages 23-24, lines 534-560

“However, the “salting-in” effect is a known phenomenon that is not likely linked with a specific physical or chemical mechanism. Quantum chemical calculation results indicated negative Gibbs free energy of water displacement for interactions between SO_4^{2-} and glyoxal monohydrate (Waxman et al., 2015). The net “salting-in” effect of 1-nitro-2-naphthol in NaF solution was interpreted by postulating hydrogen bonding (Almeida et al., 1983). A direct binding of cations to ether oxygens was proposed to be responsible for the increased solubility of water-soluble polymers (Sadeghi and Jahani, 2012). Due to the moderate correlations and negative intercepts in Figures 4 and S5, the gap between $K_{H,e}^t$ and $K_{H,w}^m$ cannot be closed by the “salting-in” effect alone. Shen et al. (2018) also obtained negative intercepts when plotting $\log(K_{H,w}^t/K_{H,e}^m)$ over c_{sulfate} for glyoxal and methylglyoxal in ambient atmosphere, and attributed this to unknown gas-particle partitioning mechanisms. Evidences showing that the thermal degradation of less volatile oligomers and organosulfates can lead to an overestimation of 2-methyltetrols by 60–188% when using a conventional GC/EI-MS method (Cui et al., 2018). To fit the gas-particle distribution of 2-methyltetrols in southeastern US, 50% of particulate 2-methyltetrols was presumed to exist in chemical forms with much lower vapor pressures by Pye et al. (2018). So, the reactive uptake and aqueous phase chemistry could be explanations for the enhanced uptake of isoprene SOA tracers. Moreover, $\log(K_{H,w}^t/K_{H,e}^m)$ values of polyol tracers also negatively correlated with the aqueous-phase concentrations of WSOC (c_{WSOC} , Figure S6), but not NH_4^+ or NO_3^- . This dependence might be associated with the “like-dissolves-like” rule, or indicate the importance of aqueous-phase heterogeneous reactions (Hennigan et al., 2009; Volkamer et al., 2009). Although several studies have estimated Henry’s law constants for a variety of polar organic compounds in pure water (e.g., polyols and polyacids; Compennolle and Müller, 2014a, b), more work is warranted to decrease the estimation uncertainty and explain their increased partitioning toward aerosol liquid water explicitly.”

As shown in the Figure below, concentrations of sulfate do not correlate with WIOC/WSOC ratios, RH, or ALWC in this work, so the condition raised in the comment was not considered.



Marcilli and Krieger (2006) found that ammonium sulfate (AS) and NaCl are “salting-out” agent for alcohols with medium hydrophilicity, including glycerol, 1,4-butanediol, 1,2-hexanediol, and PEG400. But all the target polyols in this study are

highly water soluble, and there is no evidence showing that AS is salting out agent for any of the polyols studied in this work. A number of studies identified a close relationship between salt concentrations of aerosol water and enhanced uptake of very polar compounds (Setschenow 1889; Almeida et al., 1983; Kroll et al., 2005; Ip et al., 2009; Kampf et al., 2013; Shen et al., 2018). The particle-phase interactions of organic compounds, water, and inorganic ions were found to increase partitioning of isoprene SOA tracers and levoglucosan into the condensed phase (Pye et al., 2018). Since none of the polyol compounds in Marcolli and Krieger (2006) was studied in this work, the moderate dependence of increased uptake on sulfate concentrations is not contrary to their findings, and we did not compare the study results between these two studies.

In the revised manuscript, we re-organized section 3.5 (*Partitioning coefficients of gas versus aqueous phases*) and rewrote most part of it.

18. Lines 445 – 448: related to the previous comment, here the authors state that the large gap between $K_{H,e}$ and $K_{H,w}$ cannot be explained by salting-in by sulfate alone. This confirms my impression that the discussion about reasons of the substantial deviations is rather speculative. The presented analyses do not support a firm conclusion about absorptive or non-absorptive partitioning. Moreover, if the effective Henry's law coefficient obtained is due to reactive uptake and/or aqueous phase chemistry, such as oligomer formation, then enhanced particle-phase fractions would be a reasonable expectation. However, a key question would then be whether such chemistry would be reversible during the quantification of the filter material, such that an oligomerized species would appear as monomers, since otherwise it should not contribute to the parent species' particle phase amount. This reviewer would appreciate some discussion about this.

Response:

Reactive uptake is very likely contributing to increased partitioning of polyols into the particle phase. In the revised manuscript, we have clarified that the “salting-in” effect is a known phenomenon that is not specifically linked with a physical or chemical mechanism (Pages 23-24, lines 534-560, See responses to *Comments 15 and 17*).

The equation defining “salting-in” effects was only used to parameterize the enhanced uptake of polyols in aerosol liquid water (Pages 22-23, lines 517-533).

“As sulfate has been identified as the major factor influencing the salting effect of carbonyl species (Kroll et al., 2005; Ip et al., 2009), Figure 4 shows modified Setschenow plots for C5-alkene triols, 2-methyltetrols, and levoglucosan, where $\log(K_{H,w}^t/K_{H,e}^m)$ values were regressed to the molality of sulfate ion in aerosol liquid water (c_{sulfate} , mol kg⁻¹ ALWC). The $\log(K_{H,w}^t/K_{H,e}^m)$ data increased faster when c_{sulfate} approached 0, and deviated from their expected behavior with increased c_{sulfate} . Kampf et al. (2013) selected a threshold c_{sulfate} of 12 mol kg⁻¹ ALWC to illustrate the deviation for chamber experiments, and attributed it to elevated viscosity and slow particle-phase reactions at high c_{sulfate} . In Figure 4, negative correlations ($p < 0.01$) are observed at $c_{\text{sulfate}} < 12$ mol kg⁻¹ ALWC, and Figure S5 exhibits significant negative correlations ($p < 0.01$) between $\log(K_{H,w}^t/K_{H,e}^m)$ and c_{sulfate} for individual polyols even without excluding the deviations at high c_{sulfate} . The K_s values of polyol tracers from Figures 4 and S5 (-0.17 – -0.037 kg mol⁻¹) are in a similar range as that of glyoxal (-0.24 – -0.04 kg mol⁻¹; Kampf et al., 2013; Shen et al., 2018; Waxman et al., 2015). These results

indicated that the shifting of gas-particle equilibrium toward the condensed phase might be partly parameterized by the equation defining “salting-in” effects.”

Moreover, we added some discussions about the influence of reactive uptake/aqueous chemistry on enhanced particle-phase concentrations. (Pages 23-24, lines 545-556)

“Evidences showing that the thermal degradation of less volatile oligomers and organosulfates can lead to an overestimation of 2-methyltetrols by 60–188% when using a conventional GC/EI-MS method (Cui et al., 2018). To fit the gas-particle distribution of 2-methyltetrols in southeastern US, 50% of particulate 2-methyltetrols was presumed to exist in chemical forms with much lower vapor pressures by Pye et al. (2018). So, the reactive uptake and aqueous phase chemistry could be explanations for the enhanced uptake of isoprene SOA tracers. Moreover, $\log(K_{H,w}^t/K_{H,e}^m)$ values of polyol tracers also negatively correlated with the aqueous-phase concentrations of WSOC (c_{WSOC} , Figure S6), but not NH_4^+ or NO_3^- . This dependence might be associated with the “like-dissolves-like” rule, or indicate the importance of aqueous-phase heterogeneous reactions (Hennigan et al., 2009; Volkamer et al., 2009).”

19. Table 1. Units of $K_{p,OM}$ should be provided; also state the temperature range for the values shown. Same for Table 2 and other such table in the Supplementary Information document.

Response:

Unites and temperatures were added in all related Tables as suggested (See responses to *Comment 7*).

20. Figure 2: Assuming a form of absorptive vapor–liquid equilibrium partitioning, the fraction in the particle phase of a semi-volatile organic will not only depend on the pure component vapor pressure but also on the aerosol mass concentration of the absorbing phase (and its composition). Therefore, it would make sense to state the aerosol mass concentration range that was used from the measurements. This would also allow for better comparison to other field measurements.

Response:

In this work, mass concentrations of OC were directly obtained from measurements, and were used for the calculation of $K_{p,OM}$. The concentration range of OC (2.24 – 16.8 $\mu\text{g m}^{-3}$) was added as suggested (now Figure 3, See responses to *Comment 11*).

Supplementary Information (SI):

21. Text S2: activity coefficients were assumed to be unity for all species in each sample. Is that a justified assumption? Consider that activity coefficients could be far from unity for compounds that are moderately polar (between the polarity of water and that of hydrophobic organics) used for characterizing the two particle phases in this work. This might contribute an order of magnitude of uncertainty for some compounds, but little

for others.

Response:

In this work, the organic phase in aerosols was assumed to behave as ideal solution, and the variability of activity coefficient as a function of PM composition was not considered. This should be one reason the variations of measurement-based $K_{p,OM}$ were poorly characterized in this work. Mean activity coefficients of isoprene SOA tracers and levoglucosan in organic-rich phases had a range of 0.42 to 2.04 based on AIOMFAC predictions (Pye et al., 2018), then assuming the activity coefficient to be unity will contribute an uncertainty far less than an order of magnitude.

We added some discussions to state the influence of variability in activity coefficients in the revised manuscript.

Pages 19-20, lines 450-464.

“Similarly, the measured average $F\%$ of isoprene SOA tracers in southeastern US and central Amazonia were higher than predictions by assuming instantaneous equilibrium between the gas phase and particulate OM only, and the agreement was improved when parameterization of solubility was included for predictions (Isaacman-VanWertz et al., 2016). But none of these two studies could reasonably predict the temporal variability of $F\%$ or $\log K_{p,OM}^m$. One possible explanation is that the activity coefficients of isoprene SOA tracers and levoglucosan deviate from unity (0.42–2.04; Pye et al., 2018) and vary with PM composition. Pye et al. (2018) re-analyzed the measurement data from Isaacman-VanWertz et al. (2016) using a thermodynamic equilibrium gas-particle partitioning model in two LLPS modes, which involved organic-inorganic interactions and estimations of activity coefficients as a function of liquid PM mixture composition. The resulting predictions captured both the average and diurnal variations of measured $F\%$ for polyol tracers, suggesting a necessity in obtaining time-resolved activity coefficients for the implementation of absorptive equilibrium partitioning model.”

22. Text S2, below Eq. (3): why is it the “subcooled” liquid vapor pressure? It would be sufficient to denote it as the liquid vapor pressure or liquid-state vapour pressure. Whether it is subcooled/supercooled at given temperature or just a “regular” liquid state does not matter.

Response:

The term “subcooled liquid vapor pressure” has been changed into “liquid-state vapor pressure”.

23. Also, given the relatively large uncertainty associated with vapor pressure estimation methods (O’Meara et al., 2014), it may be advised to compare those values to predictions from other methods (e.g. using the UManSysProp online tools). Uncertainties in pure-component vapor pressures could contribute more than one order of magnitude of uncertainty to K_p estimates.

Response:

Thanks. Pure-compound vapor pressures were estimated using a variety estimation tools (EPI, Evaporation, SPARC, and SIMPOL; Table S4), and the resulting $K_{p,OM}$ predictions were compared with measurement-based values in Table 1 of the revised manuscript (See responses to *Comments 7*). Descriptions of the results and associated discussions were rewritten (Section 3.4).

Pages 18-19, lines 420-464

“To understand if particulate OM is the only absorbing phase in aerosols for polyol tracers in Nanjing, the absorptive partitioning coefficients of gas vs. organic phases were calculated based on measurement results ($K_{p,OM}^m$) for predefined *Cases 1-3* and predicted theoretically ($K_{p,OM}^t$) using eq. 3 and vapor pressures listed in Table S4. In Table 1, $K_{p,OM}^t$ ranges of isoprene SOA tracers, levoglucosan, and meso-erythritol are within two orders of magnitude, while those of monosaccharides and mannitol are larger ($> 10^3$). When particulate OM was assumed as the only absorbing phase (*Case 1*), the average $K_{p,OM}^m$ of isoprene SOA tracers, levoglucosan, and meso-erythritol were more than 10 times greater than most of their $K_{p,OM}^t$ (Table 1), and this difference was not likely susceptible to measurement uncertainties. As shown in Table S5, the average relative uncertainties of measurement-based partitioning coefficients are all $< 50\%$, leading to an uncertainty of $\log K_{p,OM}^m$ less than ± 0.30 . Comparable or even greater (up to 10^5) gap between $K_{p,OM}^m$ and $K_{p,OM}^t$ has been observed for carbonyls in a number of laboratory and field studies (Healy et al., 2008; Zhao et al., 2013; Shen et al., 2018), which could be ascribed to reactive uptake (e.g., hydration, oligomerization, and esterification) of organic gases onto condensed phase (Galloway et al., 2009). Oligomers, sulfate and nitrate esters of 2-methyltetrols can be formed in the aerosol phase (Surratt et al., 2010; Lin et al., 2014), and their decomposition and hydrolysis during filter analysis will lead to an overestimation of particle-phase concentrations (Lin et al., 2013; Cui et al., 2018). However, the occurrence of oligomers, sulfate or nitrate esters of levoglucosan was not ever reported in ambient aerosols, although it can be readily oxidized by $\bullet OH$ in the aqueous phase of atmospheric particles (Hennigan et al., 2010; Hoffmann et al., 2010).

When solubility in aerosol liquid water was considered by assuming a LLPS in ambient aerosols, and whenever WSOM and WIOM partitioned into separate (*Case 2*) or single (*Case 3*) liquid phases, the average $\log K_{p,OM}^m$ of the above mentioned compounds became much closer to or even lay within the range (e.g., levoglucosan) of $\log K_{p,OM}^t$ (Table 1). These results indicated that the aerosol liquid water ($21.3 \pm 24.2 \mu g m^{-3}$; Table S1) is also an important absorbing phase of ambient polyol tracers in Nanjing. Similarly, the measured average $F\%$ of isoprene SOA tracers in southeastern US and central Amazonia were higher than predictions by assuming instantaneous equilibrium between the gas phase and particulate OM only, and the agreement was improved when parameterization of solubility was included for predictions (Isaacman-VanWertz et al., 2016). But none of these two studies could reasonably predict the temporal variability of $F\%$ or $\log K_{p,OM}^m$. One possible explanation is that the activity coefficients of isoprene SOA tracers and levoglucosan deviate from unity (0.42–2.04; Pye et al., 2018) and vary with PM composition. Pye et al. (2018) re-analyzed the measurement data from Isaacman-VanWertz et al. (2016) using a thermodynamic equilibrium gas-particle partitioning model in two LLPS modes, which involved organic-inorganic interactions and estimations of activity coefficients as a function of liquid PM mixture composition. The resulting predictions captured both the average and diurnal variations of measured $F\%$ for polyol tracers, suggesting a necessity in obtaining time-resolved activity coefficients for the implementation of absorptive equilibrium partitioning model.”

24. Figure S5: Please state the vapor pressure units and reference temperature used for the stated Log (p0L) values (one should not have to go back to the text to search for these).

Response:

The vapor pressure unit (atm) and reference temperature (25°C) were added in the figure caption of Figure S5 (now Figure S4).

“Figure S4. Temporal variations of gas-phase concentrations and particle-phase fractions ($F\%$) of polyol tracers. p°_L : Liquid-state vapor pressure (atm, EPI estimates) at 25 °C.”

References

- Almeida, M. B., Alvarez, A. M., Miguel, E. M. D., and Hoyo, E. S. D.: Setchenow coefficients for naphthols by distribution method, *Can. J. Chem.*, 61, 244-248, 10.1139/v83-043, 1983.
- Cai, J., Zeng, X., Zhi, G., Gligorovski, S., Sheng, G., Yu, Z., Wang, X., and Peng, P.: Molecular composition and photochemical evolution of water-soluble organic carbon (WSOC) extracted from field biomass burning aerosols using high-resolution mass spectrometry, *Atmos. Chem. Phys.*, 20, 6115-6128, 10.5194/acp-20-6115-2020, 2020.
- Hilal, S. H., Ayyampalayam, S. N., and Carreira, L. A.: Air–liquid partition coefficient for a diverse set of organic compounds: Henry’s law constant in water and hexadecane, *Environ. Sci. Technol.*, 42, 9231-9236, 10.1021/es8005783, 2008.
- Ip, H. S. S., Huang, X. H. H., and Yu, J. Z.: Effective Henry's law constants of glyoxal, glyoxylic acid, and glycolic acid, *Geophys. Res. Lett.*, 36, L01802, <https://doi.org/10.1029/2008GL036212>, 2009.
- Isaacman-VanWertz, G., Yee, L. D., Kreisberg, N. M., Wernis, R., Moss, J. A., Hering, S. V., de Sá, S. S., Martin, S. T., Alexander, M. L., Palm, B. B., Hu, W., Campuzano-Jost, P., Day, D. A., Jimenez, J. L., Riva, M., Surratt, J. D., Viegas, J., Manzi, A., Edgerton, E., Baumann, K., Souza, R., Artaxo, P., and Goldstein, A. H.: Ambient gas-particle partitioning of tracers for biogenic oxidation, *Environ. Sci. Technol.*, 50, 9952-9962, 10.1021/acs.est.6b01674, 2016.
- Kampf, C. J., Waxman, E. M., Slowik, J. G., Dommen, J., Pfaffenberger, L., Praplan, A. P., Prévôt, A. S. H., Baltensperger, U., Hoffmann, T., and Volkamer, R.: Effective Henry’s law partitioning and the salting constant of glyoxal in aerosols containing sulfate, *Environ. Sci. Technol.*, 47, 4236-4244, 10.1021/es400083d, 2013.
- Kroll, J. H., Ng, N. L., Murphy, S. M., Varutbangkul, V., Flagan, R. C., and Seinfeld, J. H.: Chamber studies of secondary organic aerosol growth by reactive uptake of simple carbonyl compounds, *J. Geophys. Res. Atmos.*, 110, D23207, <https://doi.org/10.1029/2005JD006004>, 2005.
- Liang, C., Pankow, J. F., Odum, J. R., and Seinfeld, J. H.: Gas/particle partitioning of semivolatile organic compounds to model inorganic, organic, and ambient smog aerosols, *Environ. Sci. Technol.*, 31, 3086-3092, 10.1021/es9702529, 1997.
- Marcilli, C., and Krieger, U. K.: Phase changes during hygroscopic cycles of mixed organic/inorganic model systems of tropospheric aerosols, *J. Phys. Chem. A*, 110, 1881-1893, 10.1021/jp0556759, 2006.
- Pankow, J. F., and Bidleman, T. F.: Interdependence of the slopes and intercepts from log-log correlations of measured gas-particle partitioning and vapor pressure—I. theory and analysis of available data, *Atmos. Environ. Part A. General Topics*, 26, 1071-1080, 10.1016/0960-1686(92)90039-n, 1992.
- Pye, H. O. T., Zuend, A., Fry, J. L., Isaacman-VanWertz, G., Capps, S. L., Appel, K. W., Foroutan, H., Xu, L., Ng, N. L., and Goldstein, A. H.: Coupling of organic and inorganic aerosol systems and the effect on gas–particle partitioning in the southeastern US, *Atmos. Chem. Phys.*, 18, 357-370, 10.5194/acp-18-357-2018, 2018.
- Setschenow, J.: Über die Konstitution der Salzlösungen auf Grund ihres Verhaltens zu Kohlensäure, *Z. Phys. Chem.*, 4U, 117-125, <https://doi.org/10.1515/zpch-1889-0409>, 1889.
- Shen, H., Chen, Z., Li, H., Qian, X., Qin, X., and Shi, W.: Gas-particle partitioning of carbonyl compounds in the ambient atmosphere, *Environ. Sci. Technol.*, 52, 10997-11006, 10.1021/acs.est.8b01882, 2018.

- Taylor, N. F., Collins, D. R., Lowenthal, D. H., McCubbin, I. B., Hallar, A. G., Samburova, V., Zielinska, B., Kumar, N., and Mazzoleni, L. R.: Hygroscopic growth of water soluble organic carbon isolated from atmospheric aerosol collected at US national parks and Storm Peak Laboratory, *Atmos. Chem. Phys.*, 17, 2555-2571, 10.5194/acp-17-2555-2017, 2017.
- Zhao, Y., Kreisberg, N. M., Worton, D. R., Isaacman, G., Weber, R. J., Liu, S., Day, D. A., Russell, L. M., Markovic, M. Z., VandenBoer, T. C., Murphy, J. G., Hering, S. V., and Goldstein, A. H.: Insights into secondary organic aerosol formation mechanisms from measured gas/particle partitioning of specific organic tracer compounds, *Environ. Sci. Technol.*, 47, 3781-3787, 10.1021/es304587x, 2013.

1 Gas-particle partitioning of polyol tracers at a suburban site in
2 the western Yangtze River Delta, China Nanjing, east China:
3 Increased partitioning to particles: Absorptive or Henry's law
4 partitioning?

5
6 Chao Qin^a, Yafeng Gou^b, Yuhang Wang^c, Yuhao Mao^b, Hong Liao^b, Qin'geng Wang^d,
7 Mingjie Xie^{b,*}

8
9
10 ^a Colleges of Resources and Environmental Sciences, Nanjing Agricultural University,
11 Nanjing 210095, China

12 ^b Collaborative Innovation Center of Atmospheric Environment and Equipment
13 Technology, Jiangsu Key Laboratory of Atmospheric Environment Monitoring and
14 Pollution Control, School of Environmental Science and Engineering, Nanjing
15 University of Information Science & Technology, 219 Ningliu Road, Nanjing 210044,
16 China

17 ^c School of Earth and Atmospheric Sciences, Georgia Institute of Technology, Atlanta,
18 GA 30332

19 ^d State Key Laboratory of Pollution Control and Resources Reuse, School of the
20 Environment, Nanjing University, Nanjing 210023, China

21
22
23
24 *Corresponding to:

25 Mingjie Xie (mingjie.xie@nuist.edu.cn; mingjie.xie@colorado.edu);

26 Tel: +86-18851903788; Fax: +86-25-58731051;

27 Mailing address: 219 Ningliu Road, Nanjing, Jiangsu, 210044, China

30

31

32 Abstract

33 Gas-particle partitioning of water-soluble organic compounds plays a significant
34 role in influencing the formation ~~and source apportionment, transport, and lifetime~~ of
35 organic aerosols in the atmosphere, but is poorly characterized. In this work, gas- and
36 particle-phase concentrations of isoprene oxidation products (C5-alkene triols and 2-
37 methylterols), levoglucosan, and sugar polyols were measured simultaneously at a
38 suburban site of the western Yangtze River Delta in east China. All target polyols
39 were primarily distributed into the particle phase (85.9—99.8%), ~~%~~. Given the
40 uncertainties in measurements and vapor pressure predictions, a dependence of ~~and~~
41 ~~their average~~ particle-phase fractions ~~were not strictly dependent~~ on vapor pressures
42 cannot be determined. To explore the impact of aerosol liquid water on gas-particle
43 partitioning of polyol tracers, three partitioning schemes (Cases 1-3) were proposed
44 based on equilibriums of gas versus organic and aqueous phases in aerosols. If
45 particulate organic matter (OM) is presumed as the only absorbing phase (Case 1), the
46 measurement-based absorptive partitioning coefficients ($K_{p,OM}^m$) of ~~Moreover, the~~
47 measurement-based partitioning coefficients ($K_{p,OM}$) of isoprene oxidation products
48 and levoglucosan were ~~more than 10^{10^2} -to- 10^4 -times~~ times larger-greater than
49 predicted values ($K_{p,OM}^t$)~~their predicted $K_{p,OM}$ -based on the equilibrium absorptive~~
50 partitioning model. The agreement between $K_{p,OM}^m$ and $K_{p,OM}^t$ was substantially
51 improved when solubility in a separate aqueous phase was included, whenever water-
52 soluble and water-insoluble OM partitioned into separate (Case 2) or single (Case 3)
53 liquid phases, suggesting that the partitioning of polyol tracers into the aqueous phase
54 in aerosols should not be ignored. These are likely attributed to the hygroscopic

55 ~~properties of polyol tracers and high aerosol liquid water (ALW) concentrations (~20~~
56 ~~$\mu\text{g m}^{-3}$) of the study location. The ~~Due to the large gaps (up to 10^7) between~~
57 ~~measurement-based effective Henry's law coefficients ($K_{\text{H,e}}^{\text{m}}$) of polyol tracers were~~
58 ~~orders of magnitude higher than and their predicted values in pure water ($K_{\text{H,w}}^{\text{t}}$),~~
59 ~~Due to the moderate correlations between $\log(K_{\text{H,e}}^{\text{m}}/K_{\text{H,w}}^{\text{t}})$ and molality of sulfate~~
60 ~~ions, the gap between $K_{\text{H,e}}^{\text{m}}$ and $K_{\text{H,w}}^{\text{t}}$ of polyol tracers could not be fully~~
61 ~~parameterized by the equation defining "salting-in" effects, and might be ascribed to~~
62 ~~mechanisms of reactive uptake, aqueous phase reaction, "like-dissolves-like"~~
63 ~~principle, etc. the gas particle partitioning of polyol tracers could not be depicted~~
64 ~~using Henry's law alone either. The regressions of $\log(K_{\text{H,w}}/K_{\text{H,e}})$ versus molality of~~
65 ~~major water soluble components in ALW indicated that sulfate ions ("salting in~~
66 ~~effect") and water soluble organic carbon can promote the partitioning of polyol~~
67 ~~tracers into the aqueous phase. These ~~These study results also partly reveals the~~
68 ~~discrepancy between results suggest a partitioning mechanism of enhanced aqueous-~~
69 ~~phase uptake for polyol tracers, which partly reveals the discrepancy between~~
70 observation and modeling of ~~secondary~~ organic aerosols.~~~~

74 **1 Introduction**

75 The water-soluble organic carbon (WSOC) ~~documented for in~~ ambient aerosols
76 can account for 20-80% of particulate organic matter (OM) based on carbon mass
77 (Saxena and Hildemann, 1996; Kondo et al., 2007). ~~Field~~ studies on the hygroscopic

78 growth and cloud condensation nucleus (CCN) activity of aerosol extracts indicated
79 that WSOC contributed significantly to aerosol hygroscopicity, and modified the
80 hydration behavior of inorganic species (e.g., sulfate, nitrate, and ammonium; Hallar
81 et al., 2013; Taylor et al., 2017). Thus, WSOC plays an important role in changing
82 radiative and cloud nucleating properties of atmospheric particles. Particulate WSOC
83 is a complex mixture of polar organic compounds containing oxygenated functional
84 groups (e.g., hydroxyl, carboxyl, and carbonyl groups), among which a list of organic
85 compounds with multiple hydroxyl (polyols) groups have been identified using gas
86 chromatography (~~GC~~)-mass spectrometry (GC-MS) and linked with specific emission
87 sources. For example, C5-alkene triols and 2-methyltetrols are isoprene oxidation
88 products (Claeys et al., 2004; Wang et al., 2005; Surratt et al., 2006); levoglucosan is
89 a typical pyrolysis product of cellulose (Simoneit et al., 1999); primary saccharides
90 (e.g., fructose and glucose) and saccharide polyols (e.g., arabitol and mannitol) are
91 commonly associated with soil microbiota and fungal spores, respectively (Simoneit
92 et al., 2004; Bauer et al., 2008).

93 To quantify the sources contributing to WSOC, concentrations of individual
94 organic tracers are often used as inputs for receptor-based modeling (Zhang et al.,
95 2009; Hu et al., 2010). Due to the influences of gas-particle partitioning on source
96 apportionment, Xie et al. (2013, 2014c) suggested the involvement of gas-phase
97 concentrations of organic markers through theoretical prediction or field measurements.
98 The equilibrium absorptive partitioning theory outlined by Pankow (1994a, b) and
99 laboratory measurements of secondary organic aerosol (SOA) yields (Odum et al.,
100 1996) have been widely applied to predict SOA formation in traditional modeling
101 studies (Heald et al., 2005; Volkamer et al., 2006; Hodzic et al., 2010). In addition to

102 absorptive partitioning to particulate OM after the formation of oxygenated organic
103 compounds in gas phase, other formation pathways (e.g., reactive uptake) have been
104 identified and are responsible for the large discrepancy between modeled and
105 observed SOA loadings ~~The large discrepancy between modeled and observed SOA~~
106 ~~loadings might be partly explained by the fact that the newly generated SOA did not~~
107 ~~undergo gas phase oxidation followed by absorptive partitioning~~ (Jang et al., 2002;
108 Kroll et al., 2005; Perraud et al., 2012). Unlike non-polar species ~~such as (e.g., n-~~
109 ~~alkanes and, polycyclic aromatic hydrocarbons (PAHs)) and alkanolic acids~~ that are
110 well simulated (Simcik et al., 1998; Xie et al., 2014a; Yatavelli et al., 2014; Isaacman-
111 VanWertz et al., 2016), ~~the absorptive partitioning model underestimated~~ particle-
112 phase concentrations of carbonyls were underestimated by several orders of
113 magnitude when particulate OM is presumed as the only absorbing phase in ideal
114 condition (Healy et al., 2008; Kampf et al., 2013; Shen et al., 2018;). Zhao et al. (2013)
115 observed a positive dependence of particle-phase pinonaldehyde on relative humidity
116 (RH, %), and inferred that aerosol water ~~avored~~ played a role in the formation of
117 pinonaldehyde in the atmosphere. However, ~~every~~ few studies have been performed
118 on the measurement of gaseous polyols (Xie et al., 2014b; Isaacman-VanWertz et al.,
119 2016), and their gas-particle partitioning were poorly understood.

120 Henry's law can ~~depict~~ describe the uptake of a compound into a liquid, highly
121 dilute solution ~~in~~ (e.g., cloud droplets) in the atmosphere (Ip et al., 2009; Compernelle
122 and Müller, 2014a). Aerosol water is also a major component of atmospheric particles,
123 and accounts for 40% by volume at 50% RH in Europe (Tsyro, 2005). But the bulk
124 aerosol solution is highly concentrated with inorganic ions and WSOC. An effective
125 Henry's law coefficient ($K_{H,e}$, mol m⁻³ atm⁻¹) can be used to account for the measured

126 partitioning between the gas phase and aerosol liquid water (Volkamer et al., 2009).
127 Both laboratory and field studies observed enhanced ~~$K_{H,e}$ effective Henry's law~~
128 ~~coefficients ($K_{H,e}^e$, mol m⁻³ atm⁻¹)~~ of carbonyl compounds with inorganic salt
129 concentrations (in mol kg⁻¹ aerosol liquid water content, ALWC; Kampf et al., 2013;
130 Waxman et al., 2015; Shen et al., 2018). This ~~termed~~ "salting-in" effect (Setschenow,
131 1889) is not mechanistically understood, and might be linked with the hydrophilic
132 interactions (e.g., hydrogen bonding) between polar organic compounds and inorganic
133 ions leading to an increase of entropy or decrease of Gibbs free energy (Almeida et al.,
134 1983; Waxman et al., 2015). Polyol tracers are highly water-soluble and their gas-
135 particle partitioning is very likely driven by the aqueous phase containing substantial
136 ionic species in ambient aerosols. In the Southeastern US, the particle-phase fraction
137 ($F\%$) of WSOC is highly dependent on RH and ALWC (Hennigan et al., 2009).

138 In the present study, polyols related to specific emission sources in gaseous and
139 particle phases were measured concurrently in northern Nanjing, China. The sampling
140 and chemical analysis were performed in a similar manner as Xie et al. (2014b), while
141 an additional step was added prior to GC-MS analysis to clean the extracts of gaseous
142 samples. To explore the roles of aerosol liquid water on gas-particle partitioning of
143 polyol tracers, three modes (Cases 1-3) were proposed based on equilibriums between
144 gas and liquid aerosol phases, and the measurement-based and predicted partitioning
145 coefficients were compared across individual cases. ~~The absorptive and Henry's law~~
146 ~~partitioning coefficients of polyol tracers were calculated based on measurements and~~
147 ~~predicted theoretically for comparison. Finally, the effects of water soluble inorganic~~
148 ~~ions and WSOC on the partitioning of atmospheric polyols were evaluated.~~ This work
149 ~~unveils~~ tends to explain the gas-particle partitioning of polyols at a suburban site in

150 eastern China, where the estimated average mass concentration of aerosol liquid water
151 is close to $20 \mu\text{g m}^{-3}$ (Yang et al., 2021). ~~The results will benefit future studies on~~
152 ~~modeling and source apportionment of organic aerosols.~~

153 **2 Methods**

154 **2.1 Field Sampling**

155 Details of the sampling information were provided in Yang et al. (2021). Briefly,
156 ambient air was sampled on the rooftop of a seven-story library building located in
157 Nanjing University of Information Science and Technology (NUIST 32.21 °N, 118.71
158 °E), a suburban site in the western Yangtze River Delta of east China. A medium
159 volume sampler (PM-PUF-300, Mingye Environmental, Gugangzhou, China)
160 equipped with a $2.5 \mu\text{m}$ cut impactor was configured to collect particulate matter with
161 aerodynamic diameter less than $2.5 \mu\text{m}$ ($\text{PM}_{2.5}$) and gaseous organic compounds at a
162 flow rate of 300 L min^{-1} . After the impactor, the sampled air flowed through a filter
163 pack containing two stacked pre-baked ($550 \text{ }^\circ\text{C}$, 4 h) quartz filters ($20.3 \text{ cm} \times 12.6 \text{ cm}$,
164 Munktell Filter AB, Sweden) and a polyurethane foam (PUF, 65 mm diameter \times 37.5
165 mm length) cartridge in series. The top quartz filter (Q_f) in the filter pack was loaded
166 with $\text{PM}_{2.5}$; gaseous organic compounds adsorbed on the backup quartz filter (Q_b) ~~was~~
167 ~~were~~ determined to evaluate sampling artifacts (~~“blow-on” and “blow-off” effects~~);
168 and the PUF cartridge was used for the sampling of gaseous polyols. ~~Xie et al. (2014b)~~
169 ~~demonstrated that using bare PUF material could capture gaseous 2-methyltetrols and~~
170 ~~levoglucosan with no excessive breakthrough ($< 33\%$).~~ Filter and PUF samples were
171 collected every sixth day during daytime (8:00 AM – 7:00 PM) and night time (7:00
172 PM – 7:00 AM next day), respectively, from 09/28/2018 to 09/28/2019. Collection
173 efficiency of gaseous polyols were examined by performing breakthrough

174 experiments using two PUF plugs during nine sampling intervals. Prior to sampling,
175 PUF adsorbents were cleaned and dried in the same way as Xie et al. (2014b). Field
176 blank filter and PUF materials were collected every 10th sample for contamination
177 adjustment. Filter and PUF samples were sealed in prebaked aluminum foil and glass
178 jars, respectively, at – 20 °C until analysis.

179 **2.2 Chemical Analysis**

180 **Bulk speciation.** The accumulated PM_{2.5} mass and bulk components including water
181 soluble ions (NH₄⁺, SO₄²⁻, NO₃⁻, Ca²⁺, Mg²⁺, and K⁺), organic (OC) and elemental
182 carbon (EC), and WSOC were measured for each filter sample. Their final
183 concentrations were determined by subtracting measurement results of Q_b from those
184 of Q_f. Concentrations of aerosol liquid water were predicted by ISORROPIA II model
185 involving using ambient temperature, RH, and concentration data of NH₄⁺, SO₄²⁻, and
186 NO₃⁻ under the metastable state. The estimated water content contributed by
187 hygroscopic WSOC was relatively small (< 1 μg m⁻³) and not accounted for in this
188 work (Text S1 of supplementary information). Table S1 lists averages and ranges of
189 ambient temperature, RH, measured PM_{2.5} components, and predicted aerosol liquid
190 water from Yang et al. (2021).

191 **Polyols analysis.** Details of the analysis method for gaseous and particulate polyols
192 were provided in supplementary information (Text ~~S1~~S2). Briefly, 1/8 of each filter
193 sample was pre-spiked with deuterated internal standard and extracted ultrasonically
194 twice for 15 min in 10–15 mL of methanol and methylene chloride mixture (1:1, v/v).
195 After filtration, rotary evaporation, N₂ blown down to dryness, and reaction with 50
196 μL of N, O-bis(trimethylsilyl)trifluoroacetamide (BSTFA) containing 1%
197 trimethylchlorosilane (TMCS) and 10 μL of pyridine, the derivatives of polyols were

198 diluted to 400 μL using pure hexane for GC-MS analysis. Pre-spiked PUF samples
199 were Soxhlet extracted using a mixture of 225 mL of methylene chloride and 25 mL
200 of methanol, followed by the same procedures of filter sample pretreatment. Prior to
201 GC-MS analysis, 50 μL of pure water was added to precipitate PUF impurities from
202 the final extract. As shown in Figure S1e, all PUF residues are kept in aqueous phase
203 at the bottom of the vial, while the derivatives of polyol tracers are supposed to be
204 retained in the top clear hexane solution. An aliquot of 2 μL of the supernatant was
205 injected for GC-MS analysis under splitless mode, and an internal standard method
206 with a six-point calibration curve (0.05–5 $\text{ng } \mu\text{L}^{-1}$) was performed to quantify polyols
207 concentrations. In this work, isoprene SOA products, including three C5-alkene triols
208 (cis-2-methyl-1,3,4-trihydroxy-1-butene, 3-methyl-2,3,4-trihydroxy-1-butene, and
209 trans-2-methyl-1,3,4-trihydroxy-1-butene; abbreviated as C5-alkene 1, 2, and 3) and
210 two 2-methyltetrols (2-methylthreitol and 2-methylerythritol), were quantified using
211 meso-erythritol; other polyols were determined using authentic standards.

212 Analytical recoveries of target polyols were obtained by adding known amounts
213 of standards to blank sampling materials (quartz filter and PUF), followed by
214 extraction and instrumental analysis identically as ambient samples. Method detection
215 limits (MDL) of individual species were estimated as three times the standard
216 deviation of their concentrations determined from six injections of the lowest
217 calibration standard. Table S2 lists recovery and MDL values of authentic standard
218 compounds. Concentrations of polyols in field blank samples were measured and
219 subtracted from air samples if necessary. To obtain appropriate gas-particle
220 distribution of polyol tracers, their field-blank corrected concentrations in filter and
221 PUF samples were adjusted by recoveries. [Final concentrations of individual polyols](#)

222 in Q_f , Q_b , and PUF samples are summarized in Table S3.

223 2.3 Data analysis

224 **Gas-particle separation and breakthrough calculation.** Polyol tracers detected in
225 Q_b samples are contributed by both gaseous adsorption (“positive artifact”) and
226 particle-phase evaporation —from Q_f samples (“negative artifact”), while—but their
227 relative contributions are unknown. ~~Previous studies rarely considered the sampling~~
228 ~~artifacts of particulate polyols.~~ Xie et al. (2014b) adjusted particle- and gas-phase
229 concentrations of levoglucosan and 2-methyltetrol based on Q_b measurements in two
230 different ways. One assumed that Q_b values were completely attributed to gaseous
231 adsorption; the other presumed equal contributions from gaseous adsorption and Q_f
232 evaporation. However, negligible difference in gas-particle distribution was observed
233 due to the small Q_b values. In Table S3, concentrations of polyol tracers on Q_b are far
234 below those on Q_f , and it would be safe to presume equal positive and negative
235 artifacts. ThusIn this study, particle-phase concentrations of polyols ~~in this study~~ were
236 represented by Q_f values, and the gas phase was calculated as the sum of Q_b and PUF
237 measurements.

238 The sampling efficiency of target polyols were evaluated by collecting and
239 analyzing tandemly installed PUF plugs during nine sampling intervals. The
240 breakthrough of each polyol was calculated as

$$241 \quad B = \frac{[\text{PUF}]_{\text{back}}}{[\text{PUF}]_{\text{front}} + [\text{PUF}]_{\text{backup}}} \times 100\% \quad (1)$$

242 where B is the breakthrough of gaseous sampling, and [PUF] represents the
243 concentration of specific compound in front or backup PUF sample. A value of 33%
244 was typically used to indicate excessive breakthrough (Peters et al., 2000; Ahrens et
245 al., 2011).

246 **Calculations of partitioning coefficients.** Here, we defined three partitioning cases
 247 to explore the influence of dissolution in aerosol liquid water on gas-particle
 248 partitioning of polyol tracers in the atmosphere. Case 1 presumes instantaneous
 249 equilibrium between the gas phase and particulate OM based on the equilibrium
 250 absorptive partitioning theory. In this case, particulate OM is assumed to be the only
 251 absorbing phase and behave as an ideal solution. Then the absorptive gas-particle
 252 partitioning coefficients ($K_{p,OM}$, $m^3 \mu g^{-1}$) were calculated from measurements ($K_{p,OM}^m$)
 253 and predicted theoretically ($K_{p,OM}^t$) as follows

$$254 \quad K_{p,OM}^m = \frac{F/M_{OM}}{A} \quad (2)$$

$$255 \quad K_{p,OM}^t = \frac{RT}{10^6 \overline{MW}_{OM} \zeta_{OM} p_L^o} \quad (3)$$

256 where M_{OM} denotes the mass concentration of absorptive organic matter ($OM = OC \times$
 257 1.6; Turpin and Lim, 2001); F ($ng\ m^{-3}$) and A ($ng\ m^{-3}$) are particulate and gaseous
 258 concentrations of individual polyols, respectively. In eq 3, R ($m^3\ atm\ K^{-1}\ mol^{-1}$) and T
 259 (K) are the ideal gas constant and ambient temperature; \overline{MW}_{OM} , average molecular
 260 weight of absorptive OM, is set at $200\ g\ mol^{-1}$ for all samples (Barsanti and Pankow,
 261 2004; Williams et al., 2010); ζ_{OM} denotes the mole fraction scale activity coefficient,
 262 and is presumed to be unity for all species in each sample; p_L^o (atm) is the vapor
 263 pressure of each pure compound, and is predicted with several estimation tools and
 264 adjusted for each sampling interval based on the average temperature (Text S3 and
 265 Table S4).he calculations of measurement and theory based absorptive partitioning
 266 coefficients ($K_{p,OM}^m$ and $K_{p,OM}^t$, $m^3\ \mu g^{-1}$) were conducted identically as those in Xie et
 267 al. (2013, 2014a, b). The equations and parameters were detailed in supplementary
 268 information (Text S2)

269 Due to the influence of mixing state and water content in aerosols, several studies
 270 modeled the gas-particle partitioning of oxygenated organic compounds by defining a
 271 liquid-liquid phase separation (LLPS) in the aerosol (Zuend and Seinfeld, 2012; Pye
 272 et al., 2018). The organic-inorganic interactions and changes of activity coefficients in
 273 aqueous mixtures were fully considered as well. As aerosol liquid water plays a
 274 significant role in the gas particle partitioning of water soluble organic compounds, In
 275 this study, we proposed an a simplified equilibrium-LLPS partitioning mechanism
 276 (Case 2) in Figure S21. First, aerosol liquid-water and water-water-insoluble OM
 277 (WIOM = OM - WSOC × 1.6) exist in two separate liquid phases (liquid-liquid phase
 278 separation), and WSOC and inorganic ions are totally dissolved in the aqueous phase.
 279 The distribution of polyol tracers between aqueous ($F_w, \text{ng m}^{-3}$) and WIOM ($F_{\text{WIOM}},$
 280 ng m^{-3}) phases is simply depicted by their octanol-water partition coefficients (K_{OW})

$$281 \quad K_{\text{OW}} = \frac{F_{\text{WIOM}}/V_{\text{WSIOM}}}{F_w/V_w} = \frac{c_{\text{WIOM}}}{c_w} \quad (24)$$

282 where V_{WIOM} and V_w are volumes (m^3) of WIOM and water in aerosols per cubic
 283 meter air; c_{WIOM} and c_w are solution concentrations (ng m^{-3}) of polyols
 284 concentrations (ng m^{-3} in solution) in WIOM-organic and aqueous phases; log K_{OW}
 285 values of target polyols were given-estimated using by the Estimation Programs
 286 Interface (EPI) Suite developed by the US Environmental Protection Agency and
 287 Syracuse Research Corporation in Table S3 (Table S4; US EPA, 2012). The density
 288 of organic matter and water (ρ_w) in aerosols are set at 1.4 and 1.0 g cm^{-3} , respectively
 289 (Isaacman-VanWertz et al., 2016; Taylor et al., 2017). Due to the high water-
 290 solubility of target polyols ($K_{\text{OW}} < 0.15$), more than 95% of their particle phase
 291 concentrations were distributed into the aqueous phase. Second, gas-phase polyol
 292 tracers are in equilibrium with hydrophobic OM and the aqueous phase, respectively;

293 following absorptive partitioning theory (eqs 1 and 2 in Text S2) and Henry's law (eq
294 3)

$$295 K_{p,WIOM}^m = \frac{F_{WIOM}/M_{WIOM}}{A} \quad (5)$$

$$296 K_{H,e}^m K_{H,e} = \frac{\frac{F_w F}{M_i}}{\frac{A}{M_i} \times R \times T \times \frac{c_{ALW}}{\rho_w}} = \frac{\rho_w \times F_w F}{A \times R \times T \times c_{ALW}} \quad (36)$$

297 where $K_{H,e}^m$ ($\text{mol m}^{-3} \text{ atm}^{-1}$) is the measurement-based effective Henry's law
298 coefficient; M_{WIOM} represents the mass concentration ($\mu\text{g m}^{-3}$) of WIOM; F and A (ng
299 m^{-3}) represent particle and gas phase concentrations of polyol tracers in ambient air;
300 M_i (g mol^{-1}) is the molecular weight of specific compound; R ($\text{m}^3 \text{ atm K}^{-1} \text{ mol}^{-1}$) and T
301 (K) are ideal gas constant and ambient temperature, respectively; c_{ALW} ($\mu\text{g m}^{-3}$) is the
302 mass concentration of aerosol liquid water predicted using ISORROPIA II model; ρ_w
303 (1 g cm^{-3}) is water density. Case 3 is generally the same as Case 2, and the only
304 difference is that water-soluble OM (WSOM) and WIOM exist in a single organic
305 phase. Here total particulate OM was used instead of WIOM to assess the distribution
306 of polyol tracers between aqueous and organic phases, and calculate partitioning
307 coefficients of gas vs. particulate organic ($K_{p,OM}^m$) and aqueous ($K_{H,e}^m$) phases. Note
308 that the polarity of particulate OM phase in Case 3 was expected to increase, then
309 using K_{OW} to calculate the distribution of polyols between organic and aqueous phases
310 might lead to underestimated $K_{p,OM}^m$ and overestimated $K_{H,e}^m$. Due to the high water-
311 solubility of target polyols ($K_{OW} < 0.15$), more than 95% of their particle phase
312 concentrations were distributed into the aqueous phase. For comparison purposes, the
313 Henry's law coefficient in pure water at 25 °C ($K_{H,w}^*$) was ~~predicted~~ estimated from
314 using EPI suite and SPARC (Hilal et al., 2008; <http://archemcalc.com/sparc-web/calc>),
315 respectively (Table S3S4), and was adjusted for each sampling interval due to the

316 changes in ambient temperature using van_'t Hoff equation (Text S3S4).

317 Uncertainty estimation. To obtain the uncertainty associated with the calculation of $F\%$
318 and partitioning coefficients ($K_{p,OM}^m$ and $K_{H,e}^m$), measurement uncertainties of polyol
319 tracers in filter and PUF samples were estimated from their recoveries and
320 breakthrough for gaseous sampling. The root sum of squares (RSS) method was
321 applied to propagate uncertainties of gas and particle-phase concentrations for $F\%$,
322 $K_{p,OM}^m$, and $K_{H,e}^m$ calculations. Details of the uncertainty estimation and propagation
323 methods were provided in Text S5, and the average relative uncertainties were
324 summarized in Table S5.

325 **3 Results and discussion**

326 **3.1 Method evaluation**

327 In our previous study, PUF/XAD-4 resin/PUF and PUF/XAD-7 resin/PUF
328 adsorbent sandwiches were tested for sampling gaseous 2-methyltetrols and
329 levoglucosan (Xie et al., 2014b). The results of breakthrough experiments suggested
330 that both the two sandwiched composites had high sampling efficiency (close to
331 100%). Moreover, individual parts of the two types of composites (top PUF, middle
332 XAD-4/XAD-7 resin, and backup PUF) were analyzed for 7 samples, and target
333 compounds were only detected in top PUF. It is therefore suitable to collect gaseous
334 2-methylterols and levoglucosan using PUF materials only~~Thus, bare PUF material is~~
335 ~~suitable for sampling gaseous 2-methyltetrols and levoglucosan.~~

336 Although PUF materials were pre-cleaned prior to sampling, a few short-chain
337 polyurethanes or impurities could be dissolved during Soxhlet extraction of target
338 compounds using the mixture of methanol and methylene chloride. These substances
339 precipitated when sample extracts were concentrated (Figure S1a, b), and re-dissolved

340 in BSTFA:TMCS/pyridine and hexane after the derivatization step (Figure S1c, d). In
341 Xie et al. (2014b), an aliquot of 2 μL of the sample extract as shown in Figure S1d
342 was injected for GC-MS analysis. ~~Due to the fact that~~ Since the dissolved PUF
343 materials did not vaporize at ~ 300 $^{\circ}\text{C}$, the GC inlet liner had to be changed for
344 cleaning every few samples. In this work, 50 μL of pure water was added to separate
345 PUF materials from polyol derivatives in hexane solution. As shown in Figure S1e, all
346 PUF residues were retained in the aqueous solution after ~~liquid-liquid~~-phase
347 separation. This pretreatment step was added for the analysis of gaseous samples to
348 save time for changing and cleaning GC inlet liners. However, the revised method did
349 not improve the recoveries of meso-erythritol and levoglucosan in PUF samples
350 (Table S2) compared to those in Xie et al. (2014b). This is because the dissolved PUF
351 materials should have an impact on the derivatization efficiency of polyol species, and
352 future work is warranted to remove dissolved PUF materials in sample extracts before
353 the derivatization step.

354 Measurement results of breakthrough samples and the resulting B values were
355 shown in Figure ~~S3S2~~. C5-alkene triols and 2-methyltetrols were mainly observed in
356 summertime, and levoglucosan was only detected in three pairs of breakthrough
357 samples. Their average B values ($< 33\%$) indicated no excessive breakthrough (Figure
358 ~~S3aS2a-c~~), but were higher than those reported by Xie et al. (2014b). This might be
359 ascribed to the greater face velocity (1.5 cm s^{-1}) for sampling gaseous polyols than
360 that (0.61 cm s^{-1}) in our previous study. Due to the limit in sample number for
361 breakthrough tests and low detection rates, the dependence of breakthrough on
362 ambient temperature or OM loadings cannot be evaluated. The breakthrough of an
363 ideal sampling method is expected to be extremely low (e.g., $< 10\%$) and have no

364 dependence on ambient temperature, OM loadings, etc. Unlike fructose which had
365 low breakthrough (Figure ~~S3d~~S2d), glucose and mannitol had comparable
366 concentrations between front and backup PUF samples for several breakthrough
367 experiments (Figure ~~S3e~~S2e, f), indicating that PUF materials are not suitable for
368 sampling gaseous glucose and mannitol. Mannose and arabitol were not detected or
369 had BDL values for breakthrough samples, and their breakthrough was not provided.
370 In the current work, concentrations of polyol tracers in filter and PUF samples were
371 all reported, but the data of mannose, glucose, arabitol, and mannitol in PUF samples
372 should be treated with caution due to high breakthrough or the lack of valid
373 breakthrough results.

374 **3.2 General description of measurement results**

375 ~~Concentrations of individual polyols in Q_f , Q_b , and PUF samples are summarized~~
376 ~~in Table S4, and their \pm Total ambient concentrations ($Q_f + Q_b + \text{PUF}$)~~ of individual
377 polyols are depicted using boxplots in Figure ~~12~~. Figure ~~S4-S3~~ presents temporal
378 variations of total and Q_f concentrations of individual polyols with daytime and night-
379 time measurements distinguished. In general, polyol tracers were predominantly
380 observed on Q_f with averages 1-3 orders of magnitude higher than those on Q_b and
381 PUF (Table S3). Levoglucosan had the highest average total concentration ($66.1 \pm$
382 71.1 ng m^{-3}), followed by fructose ($15.0 \pm 62.9 \text{ ng m}^{-3}$) and mannose ($14.3 \pm 31.3 \text{ ng}$
383 m^{-3}). C5-alkene triols and 2-methyltetrols are formed from isoprene epoxydiols
384 (IEPOX) under low NO_x conditions (Surratt et al., 2010). All the five species on Q_b
385 were more frequently detected and had average concentrations 2-20 times higher than
386 those in PUF samples. While in Xie et al. (2014b), the sum of 2-methyltetrols in Q_b
387 and adsorbent samples were up to 2.7 times higher than those on Q_f in summer

388 Denver, so isoprene products are not similarly distributed between gas and aerosol
389 phases across different regions. Moreover, isoprene-derived polyols exhibited
390 prominent elevations in summer (Figure S4aS3a-e), and their daytime concentrations
391 ($2.02 \pm 3.73 - 10.5 \pm 29.3 \text{ ng m}^{-3}$) were only slightly higher than those during night-
392 time ($1.63 \pm 4.40 - 9.65 \pm 32.7 \text{ ng m}^{-3}$). Previous field studies observed strong diurnal
393 variations of isoprene SOA tracers with peak concentrations from afternoon till
394 midnight (Fu and Kawamura, 2011; Isaacman-VanWertz et al., 2016). Fu and
395 Kawamura (2011) investigated diurnal variations of polar organic tracers at a forest
396 site in summer by sampling aerosol particles every 4 h. They found that isoprene-
397 derived SOA tracers maximized from later afternoon to early evening. Although no
398 IEPOX will be generated from the oxidation of isoprene by $\bullet\text{OH}$ and $\text{HO}_2\bullet$ after
399 sunset, the formations of C5-alkene triols and 2-methyltetrols might continue until
400 pre-existing IEPOX is exhausted. In this work, neither the daytime (8:00 AM–7:00
401 PM) or night-time (7:00 AM–7:00 AM next day) sample covered the whole period
402 when isoprene SOA tracers had peak concentrations, and the strong diurnal variations
403 of C5-alkene triols and 2-methyltetrols were not captured. This explains the
404 insignificant ($p > 0.05$) day night differences of C5-alkene triols and 2-methyltetrols
405 in this work.

406 Levoglucosan was more frequently detected but far less concentrated in PUF than
407 in Q_b samples. Its total concentrations were comparable to those in urban Denver
408 (average $65.3 \pm 96.8 \text{ ng m}^{-3}$, range 2.48 – 478 ng m^{-3}), where an average of ~20%
409 partitioned into the gas phase (Xie et al., 2014b). Due to the enhanced biomass
410 burning activities in cold periods for domestic heating at night, levoglucosan showed
411 a clear seasonal pattern (winter maxima and summer minima) and significant ($p =$

412 0.03) higher concentrations during night-time (Figure [S4fS3f](#)). Sugars and sugar
413 alcohols are commonly linked with soil/dust resuspension and associated microbial
414 activities (Simoneit et al., 2004). They were frequently detected in Q_b samples with
415 comparable averages and ranges as those in PUF samples (Table [S4S3](#)). Total
416 concentrations of fructose and glucose were strongly ($r = 0.98$) correlated peaking in
417 middle spring (April 2019, Figure [S4hS3h](#), j), when Ca^{2+} on Q_f also reached its
418 maxima of the year (Yang et al., 2021), indicating an influence from soil/dust
419 resuspension. Arabitol and mannitol had identical seasonal pattern ($r = 0.89$) with
420 elevated total concentrations from May to October (Figure [S4iS3i](#), m), which might
421 be attributed to ~~the~~ high levels of vegetation during growing seasons and autumn
422 decomposition decomposing (Burshtein et al., 2011). Multiple peaks of mannose
423 concentrations were observed from spring to autumn, suggesting a variety of
424 contributing sources (e.g., microbial activity, vegetation). Xylitol is likely derived
425 from biomass burning in northern Nanjing due to its strong correlation ($r = 0.89$) with
426 levoglucosan.

427 **3.3 Gas- and particle-phase distributions and absorptive partitioning coefficient**

428 ~~Q_b measurements were often used to assess positive sampling artifacts of~~
429 ~~particulate OC (Chow et al., 2010; Subramanian et al., 2004), but rarely for particle-~~
430 ~~phase organic markers. In this study~~ As mentioned in sections 2.3, concentrations of
431 particulate polyols were obtained directly from Q_f measurements, and the gas phase
432 was calculated as the sum of Q_b and PUF values. Figure [S5-S4](#) shows the time series
433 of gas-phase concentrations and particle-phase fractions ($F\%$) of individual polyol
434 tracers. The average $F\%$ values of measured species are ~~linearly regressed~~ plotted
435 against the logarithms of their ~~subcooled~~ liquid-state vapor pressures at 25 °C ($p^{0,*}_L$)

436 in Figure 23. Gas-phase C5-alkene triols and 2-methyltetrols had maximum
437 concentrations in summer and significant ($p < 0.05$) day-night variations (Figure S4),
438 while other polyols had extremely low concentrations in the gas phase with $F\%$
439 (average \pm standard deviation) ranging from $94.2 \pm 8.02 - 99.8 \pm 1.21\%$. The
440 average $F\%$ values of 2-methyltetrols ($87.5 \pm 10.6\%$) and levoglucosan ($99.8 \pm 1.21\%$)
441 were greater than those in urban Denver (50–80%; Xie et al., 2014b), where the
442 average sampling temperature (12.5 ± 10.1 °C) was much lower. Thus, the changes in
443 vapor pressures with the ambient temperature and/or particulate OM loadings might
444 not be the main factors driving gas-particle partitioning of polyol tracers in Nanjing.
445 In Figure 3, the average $F\%$ uncertainties (6.16–31.2%) of monosaccharides (e.g.,
446 fructose) and sugar alcohols (e.g., mannitol) were larger than those of isoprene SOA
447 tracers and levoglucosan (3.33–7.24%) due to their low and variable recoveries (Table
448 S2) and excessive breakthrough (Figure S2). However, the estimated uncertainties of
449 $F\%$ for less volatile polyols ($p^{0,*}_L < \sim 10^{-10}$ atm) were not physically meaningful, as
450 more than 95% of these compounds existed in the particle phase. Considering the
451 uncertainties in $F\%$ and $\log p^{0,*}_L$ and high average $F\%$ ($> 85\%$) of target polyol
452 tracers, a dependence of $F\%$ on the vapor pressure could not be determined, and the
453 seasonality and day-night difference ($p > 0.05$) of $F\%$ were obscured. Unlike non-
454 polar organic tracers (e.g., n -alkanes and PAHs), some polyols (e.g., 2-methyltetrols
455 and levoglucosan) data did not follow the linear regression line of $F\%$ versus $\log p^{0,*}_L$.
456 Although gas-phase C5-alkene triols and 2-methyltetrols were majorly observed in
457 summer with significant ($p < 0.05$) day-night variations, their $F\%$ values did not show
458 seasonality or day-night difference ($p = 0.18-0.73$). Other polyols had extremely low
459 concentrations in the gas phase with average $F\%$ ranging from $94.2 \pm 8.02 - 99.8 \pm$

460 1.21%. The average $F\%$ values of 2-methyltetrols ($87.5 \pm 10.6\%$) and levoglucosan
461 ($99.8 \pm 1.21\%$) were greater than those in urban Denver (50–80%; Xie et al., 2014b),
462 where the average sampling temperature ($12.5 \pm 10.1\text{ }^\circ\text{C}$) was much lower. Thus, the
463 changes in vapor pressures with the ambient temperature might not be the main factor
464 driving gas-particle partitioning of polyol tracers in northern Nanjing.

465 3.4 Partitioning coefficients of gas versus organic phases

466 To understand if particulate OM is the only absorbing phase in aerosols for
467 polyol tracers in Nanjing, the absorptive partitioning coefficients of gas vs. organic
468 phases were calculated based on measurement results ($K_{p,OM}^m$) for predefined Cases
469 1-3 and predicted theoretically ($K_{p,OM}^t$) using eq. 3 and vapor pressures listed in Table
470 S4. In Table 1, $K_{p,OM}^t$ ranges of isoprene SOA tracers, levoglucosan, and meso-
471 erythritol are within two orders of magnitude, while those of monosaccharides and
472 mannitol are larger ($> 10^3$). understand if traditional absorptive partitioning theory
473 could be applied to predict the gas-particle partitioning of polyol tracers in northern
474 Nanjing. When particulate OM was assumed as the only absorbing phase (Case 1),
475 Table 1 compares $\log K_{p,OM}^m$ and $\log K_{p,OM}^t$ of individual compounds. The the
476 average $K_{p,OM}^m$ values of isoprene SOA tracers, levoglucosan, and meso-erythritol
477 were 10^2 to 10^3 more than 10 times larger-greater than most of their corresponding
478 $K_{p,OM}^t$ (Table 1), and this difference was not likely susceptible to measurement
479 uncertainties. As shown in Table S5, the average relative uncertainties of
480 measurement-based partitioning coefficients are all $<50\%$, leading to an uncertainty
481 of $\log K_{p,OM}^m$ less than ± 0.30 . Comparable or even greater (up to 10^5) gap between
482 $K_{p,OM}^m$ and $K_{p,OM}^t$ has been observed for carbonyls in a number of laboratory and field
483 studies (Healy et al., 2008; Zhao et al., 2013; Shen et al., 2018), which could be

484 ascribed to reactive uptake (e.g., hydration, oligomerization, and esterification) of
485 organic gases onto condensed phase (Galloway et al., 2009). Oligomers, sulfate and
486 nitrate esters of 2-methyltetrols can be formed in the aerosol phase (Surratt et al.,
487 2010; Lin et al., 2014), and their decomposition and hydrolysis during filter analysis
488 will lead to an overestimation of particle-phase concentrations (Lin et al., 2013; Cui et
489 al., 2018). ~~but these products were not expected to dominate particle phase~~
490 ~~concentrations of 2-methyltetrols (Lin et al., 2013; Xie et al., 2014b).~~ However,
491 ~~Although levoglucosan can be readily oxidized by •OH in the aqueous phase of~~
492 ~~atmospheric particles (Hennigan et al., 2010; Hoffmann et al., 2010),~~ the occurrence
493 of ~~its~~ oligomers, sulfate or nitrate esters of levoglucosan was not ever reported in
494 ambient aerosols. A, although levoglucosan it can be readily oxidized by •OH in the
495 aqueous phase of atmospheric particles (Hennigan et al., 2010; Hoffmann et al.,
496 2010).

497 When solubility in aerosol liquid water was considered by assuming a LLPS in
498 ambient aerosols, and whenever WSOM and WIOM partitioned into separate (Case 2)
499 or single (Case 3) liquid phases, the average log $K_{p,OM}^m$ of the above mentioned
500 compounds became much closer to or even lay within the range (e.g., levoglucosan)
501 of log $K_{p,OM}^l$ (Table 1). These results indicated that the aerosol liquid water ($21.3 \pm$
502 $24.2 \mu\text{g m}^{-3}$; Table S1) is also an important absorbing phase of ambient polyol tracers
503 in Nanjing. Xie et al. (2014b) found that the gas particle partitioning of 2-
504 methyltetrols and levoglucosan in urban Denver were highly dependent on the
505 variations in ambient temperature and absorbing organic matter (M_{OM}). Similarly,
506 While in southeastern US, the particle-phase fractions measured average $F\%$ of
507 isoprene SOA tracers in southeastern US and central Amazonia were generally higher

508 than predictions by assuming instantaneous equilibrium between the gas phase and
509 particulate OM only, and the agreement was improved when parameterization of
510 solubility was included for predictions (Isaacman-VanWertz et al., 2016). But none of
511 these two studies could reasonably predict the temporal variability of $F\%$ or log
512 $K_{p,OM}^m$. One possible explanation is that the activity coefficients of isoprene SOA
513 tracers and levoglucosan deviate from unity (0.42–2.04; Pye et al., 2018) and vary
514 with PM composition. Pye et al. (2018) re-analyzed the measurement data from
515 Isaacman-VanWertz et al. (2016) using a thermodynamic equilibrium gas-particle
516 partitioning model in two LLPS modes, which involved organic-inorganic interactions
517 and estimations of activity coefficients as a function of liquid PM mixture
518 composition. The resulting predictions captured both the average and diurnal
519 variations of measured $F\%$ for polyol tracers, suggesting a necessity in obtaining
520 time-resolved activity coefficients for the implementation of absorptive equilibrium
521 partitioning model.

522 ~~based on absorptive partitioning model (Isaacman-VanWertz et al., 2016). This~~
523 ~~discrepancy might be related to the spatial heterogeneity of ALWC, which is expected~~
524 ~~to control the gas-particle partitioning of water-soluble organic matter in the Eastern~~
525 ~~US (Carlton and Turpin, 2013). In this study, the large difference between $K_{p,OM}^m$ and~~
526 ~~$K_{p,OM}^t$ indicated that some mechanisms other than absorptive partitioning (e.g.,~~
527 ~~Henry's law partitioning) should be involved to predict the gas-particle partitioning~~
528 ~~of polyol tracers in northern Nanjing, where the ambient particles contained~~
529 ~~substantial liquid water ($21.3 \pm 24.2 \mu\text{g m}^{-3}$; Table S1).~~

530 Unlike isoprene SOA tracers and levoglucosan, the average $K_{p,OM}^t$ values of
531 monosaccharides (fructose, mannose, and glucose) and sugar alcohols (xylitol,

532 arabitol, and mannitol) were ~~up to 10^3 times~~ orders of magnitude larger than their
533 $K_{p,OM}^m$ for Cases 2 and 3 (Table 1). This is probably caused by the overestimation of
534 gas-phase concentrations of sugar polyols. The organic matter on Q_b is mainly
535 composed of volatile and semi-volatile organic compounds. If the concentrations of
536 organic compounds on Q_b were comparable or higher than those on Q_f , their Q_f values
537 should be dominated by ~~positive artifact~~ positive artifact. As the vapor pressure
538 decreases, the evaporation loss from Q_f samples becomes non-negligible. Note that
539 the magnitude of negative artifacts is unknown and very difficult to assess, and the
540 vapor pressures of monosaccharides and sugar alcohols are ~~mostly~~ far below 10^{-10}
541 atm (Table S4), their concentrations in Q_b and even PUF samples might contain more
542 contributions from negative artifacts than isoprene SOA tracers and levoglucosan.
543 ~~Considering that~~ As low-volatile sugar polyols had lower and less stable recoveries
544 (Table S2) and greater breakthrough (Figure ~~S3e~~ S2e, f), caution is warranted in
545 analyzing their $K_{p,OM}^m$ values obtained in this study.

546 ~~Figure S2 presumes that gas-phase polyols are in equilibrium with WIOM and the~~
547 ~~aqueous phase, respectively. Then concentrations of WIOM [$1.4 \times (OC - WSOC)$] was~~
548 ~~used to adjust the calculation of absorptive partitioning coefficients ($K_{p,WIOM}^m$) based~~
549 ~~on eq 1 in supplementary information. In comparison to $\log K_{p,OM}^m$, the average \log~~
550 ~~$K_{p,WIOM}^m$ values of isoprene SOA tracers and levoglucosan were much closer to~~
551 ~~average $\log K_{p,OM}^t$ (Tables 1 and S5), supporting that the aerosol liquid water should~~
552 ~~have significant impacts on gas-particle partitioning of polyol tracers.~~

553 3.4-5 Partitioning coefficients of gas versus aqueous phases ~~Effective Henry's law~~ 554 ~~coefficient~~

555 Table 2 lists the statistics of measurement-based $\log K_{H,e}$ and predicted $\log K_{H,w}$.

556 The predicted Henry's law coefficients in pure water ($K_{H,w}^t$, mol m⁻³ atm⁻¹) from EPI
557 and SPARC estimates differed by several orders of magnitude, but literature values of
558 isoprene SOA and levoglucosan were closer to EPI estimates (Table S4). If SPARC
559 $K_{H,w}^*$ values were used, the average log $K_{H,e}^m$ of most polyol tracers would be lower
560 than their average log $K_{H,w}^t$ (Table 2), indicating that the aqueous phase of ambient
561 aerosol is less hospitable to polyol tracers than pure water. This is in conflict with the
562 fact that the interactions of organic compounds, water, and inorganic ions in aerosols
563 will increase the partitioning of highly oxygenated compounds (O:C ≥ 0.6; e.g.,
564 isoprene SOA tracers and levoglucosan) into the particle phase (Pye et al., 2018).
565 Several studies identified a close relationship between salt concentrations of aerosol
566 water and enhanced uptake of very polar compounds (Almeida et al., 1983; Kroll et
567 al., 2005; Ip et al., 2009; Kampf et al., 2013; Shen et al., 2018). Thus, log $K_{H,w}^t$ values
568 of EPI estimates were used for further data analysis.

569 In Table 2, the average $K_{H,w}^t/K_{H,w}$ values of isoprene SOA tracers, levoglucosan,
570 and meso-erythritols based on EPI estimations were 2–6 orders of magnitude 10^2 to 10^6
571 lower than their corresponding average $K_{H,e}^m/K_{H,e}$. Log $K_{H,e}^m$ values of Cases 2 and 3
572 had ignorable difference and were not presented separately, indicating that the
573 ambient atmosphere in northern Nanjing favored the condensation of these polyols.
574 Other polyol compounds exhibited less difference between log $K_{H,e}^m/K_{H,e}$ and log
575 $K_{H,w}^t/K_{H,w}$, which was very likely caused by the overestimation of their gas-phase
576 concentrations. The average $K_{H,e}^m$ values of polyol tracers (10^{13} – 10^{15} mol m⁻³ atm⁻¹) in
577 this study were several orders of magnitude larger than those of carbonyls derived
578 from ambient measurements (10^{10} – 10^{12} mol m⁻³ atm⁻¹; Shen et al., 2018) and chamber
579 simulations ($\sim 10^{11}$ mol m⁻³ atm⁻¹; Kroll et al., 2005; Volkamer et al., 2006; Galloway

580 et al., 2009). This is because low molecular weight carbonyls (e.g., glyoxal) are much
581 more volatile ($p^{0,*}_L > 10^{-2}$ atm) than our target polyols (Table S4). According to
582 existing studies, the minimum concentrations of gas-phase glyoxal and methylglyoxal
583 in Chinese cities ($\sim 0.1 \mu\text{g m}^{-3}$) are magnitudes higher than the averages of polyol
584 tracers in this work, while their particle-phase concentrations are of the same
585 magnitude (Shen et al., 2018; Liu et al., 2020).

586 A number of previous studies observed enhanced $K_{H,e}$ of carbonyls with salt
587 concentrations in aqueous solution (Ip et al., 2009; Kampf et al., 2013; Waxman et al.,
588 2015; Shen et al., 2018), and described this “salting-in” effect using

$$589 \text{Log} \left(\frac{K_{H,w}}{K_{H,e}} \right) = K_s c_{\text{salt}} \quad (47)$$

590 where K_s (kg mol^{-1}) is the salting constant, and c_{salt} is the aqueous-phase concentration
591 of salt in mol kg^{-1} ALWC. This equation is originally defined in Setschenow (1889)
592 by plotting $\log (K_{H,w}/K_{H,e})$ versus the total salt concentration (mol L^{-1}).

593 As sulfate has been identified as the major factor influencing the salting effect of
594 carbonyl species (Kroll et al., 2005; Ip et al., 2009), Figure 3–4 shows modified
595 Setschenow plots for C5-alkene triols, 2-methyltetrols, and levoglucosan, where \log
596 ($K^1_{H,w}/K^m_{H,e}$) values were regressed to the molality of sulfate ion in aerosol liquid
597 water (c_{sulfate} , mol kg^{-1} ALWC). ~~However, the $\log (K^1_{H,w}/K^m_{H,e})$ data~~
598 increased faster when c_{sulfate} approached 0, and deviated from their expected behavior
599 with in the modified Setschenow plot increased c_{sulfate} . Kampf et al. (2013) selected a
600 threshold c_{sulfate} of 12 mol kg^{-1} ALWC to illustrate the deviation for chamber
601 experiments, and attributed it to elevated viscosity and slow particle-phase reactions
602 at high c_{sulfate} . at $c_{\text{sulfate}} > 12 \text{ mol kg}^{-1}$ ALWC, which was also observed for glyoxal
603 (Kampf et al., 2013). This might be because the ambient particles did not undergo

604 liquid-liquid phase separation at $c_{\text{sulfate}} > 12 \text{ mol kg}^{-1}$ ALWC, when the average RH
605 ($51.5 \pm 15.4\%$) was lower than the lowest deliquescence RH (61.8%) of major
606 inorganic salts (e.g., NH_4NO_3 , $(\text{NH}_4)_2\text{SO}_4$) in ambient aerosols (Seinfeld and Pandis,
607 2016), and the corresponding average concentration of aerosol liquid water was only
608 $5.31 \pm 4.05 \mu\text{g m}^{-3}$. In Figure 34, negative correlations ($p < 0.01$) are observed at
609 $c_{\text{sulfate}} < 12 \text{ mol kg}^{-1}$ ALWC, and Figure S5 exhibits significant negative correlations
610 ($p < 0.01$) between $\log(K_{\text{H,w}}^{\text{t}}/K_{\text{H,e}}^{\text{m}})$ and c_{sulfate} for individual polyols even without
611 excluding the deviations at high c_{sulfate} . The K_{s} values of polyol tracers from Figures 4
612 and S5 (-0.17 – $-0.037 \text{ kg mol}^{-1}$) are in a similar range as that of glyoxal (-0.24 – -0.04
613 kg mol^{-1} ; Kampf et al., 2013; Shen et al., 2018; Waxman et al., 2015). These results
614 indicated that the shifting of gas-particle equilibrium toward the condensed phase
615 might be partly parameterized by the equation defining “salting-in” effects.

616 and the K_{s} values range from -0.17 to $-0.15 \text{ kg mol}^{-1}$. Figure S6 shows the
617 regressions between $\log(K_{\text{H,w}}/K_{\text{H,e}})$ of individual polyols and c_{sulfate} without
618 considering the deviations at high c_{sulfate} , and nearly all species exhibit significant
619 negative correlations ($p < 0.01$). These results indicated the “salting in” effects for
620 polyol tracers in northern Nanjing, and to our knowledge the present study is the first
621 to calculate their $K_{\text{H,e}}$ and K_{s} . Although several studies have estimated Henry’s law
622 constants for a variety of polar organic compounds in pure water (e.g., polyols and
623 polyacids; Compernelle and Müller, 2014a, b), salting effects should be considered in
624 describing their gas-particle partitioning in the ambient atmosphere.

625 The average $K_{\text{H,e}}$ values of polyol tracers (10^{13} – $10^{15} \text{ mol m}^{-3} \text{ atm}^{-1}$) in this study
626 were several orders of magnitude larger than those of carbonyls derived from ambient
627 measurements (10^{10} – $10^{12} \text{ mol m}^{-3} \text{ atm}^{-1}$; Shen et al., 2018) and chamber simulations

628 ($\sim 10^{11}$ mol m⁻³ atm⁻¹; Kroll et al., 2005; Volkamer et al., 2006; Galloway et al., 2009).
629 This is because low molecular weight carbonyls (e.g., glyoxal) are much more volatile
630 ($p^{\text{sat}}_{\text{L}} > 10^2$ atm) than our target polyols (Table S3). According to existing studies, the
631 minimum concentrations of gas phase glyoxal and methylglyoxal in Chinese cities
632 (~ 0.1 $\mu\text{g m}^{-3}$; Liu et al., 2020) are magnitudes higher than the averages of polyol
633 tracers in this work, while their particle phase concentrations are of the same
634 magnitude. The K_s values of polyol tracers from Figures 3 and S5 (0.17–0.037 kg
635 mol⁻¹) are in a similar range as that of glyoxal (0.24–0.04 kg mol⁻¹; Kampf et al.,
636 2013; Shen et al., 2018; Waxman et al., 2015), indicating that the uptake of different
637 water soluble organic compounds might be enhanced by sulfate in a similar manner.
638 However, the mechanisms of “salting in” effects are not fully understood. Kampf et al.
639 (2013) inferred that the enhanced uptake of glyoxal was accompanied by chemical
640 reactions in the aqueous phase (e.g., hydration and oligomerization), and the
641 interactions between SO₄²⁻ and glyoxal monohydrate had negative Gibbs free energy
642 of water displacement (Waxman et al., 2015). The net “salting in” effect of 1-nitro-2-
643 naphthol in NaF solution was interpreted by postulating hydrogen bonding (Almeida
644 et al., 1983). A direct binding of cations to ether oxygens was proposed to be
645 responsible for the increased solubility of water soluble polymers (Sadeghi and Jahani,
646 2012).

647 Due to the complexity of PM composition However, the “salting-in” effect is a
648 known phenomenon that is not likely linked with a specific physical or chemical
649 mechanism. Quantum chemical calculation results indicated negative Gibbs free
650 energy of water displacement for interactions between SO₄²⁻ and glyoxal monohydrate
651 (Waxman et al., 2015). The net “salting-in” effect of 1-nitro-2-naphthol in NaF

652 solution was interpreted by postulating hydrogen bonding (Almeida et al., 1983). A
653 direct binding of cations to ether oxygens was proposed to be responsible for the
654 increased solubility of water-soluble polymers (Sadeghi and Jahani, 2012). Due to the
655 moderate correlations and negative intercepts in Figures 4 and S5, the large gap
656 between $K_{H,e}^l$ and $K_{H,w}^m$ cannot be closed by the “salting-in” effect alone, which is
657 supported by the Shen et al. (2018) also obtained negative intercepts when plotting
658 $\log (K_{H,w}^l/K_{H,e}^m)$ over c_{sulfate} for glyoxal and methylglyoxal in ambient atmosphere,
659 and attributed this to unknown gas-particle partitioning mechanisms. Evidences
660 showing that the thermal degradation of less volatile oligomers and organosulfates can
661 lead to an overestimation of 2-methyltetrols by 60–188% when using a conventional
662 GC/EI-MS method (Cui et al., 2018). To fit the gas-particle distribution of 2-
663 methyltetrols in southeastern US, 50% of particulate 2-methyltetrols was presumed to
664 exist in chemical forms with much lower vapor pressures by Pye et al. (2018). So, the
665 reactive uptake and aqueous phase chemistry could be explanations for the enhanced
666 uptake of isoprene SOA tracers. Moreover, negative intercepts of linear regressions
667 in Figure 3. As shown in Figures S7, $\log (K_{H,w}^l/K_{H,e}^m)(K_{H,w}/K_{H,e})$ values of polyol
668 tracers also negatively correlated with the aqueous-phase concentrations of WSOC
669 (c_{wsoc} , Figure S6), but not NH_4^+ or NO_3^- . given that the plots are more scattered at
670 high c_{wsoc} . This dependence might be partly explained using associated with the
671 “like-dissolves-like” rule, and indicate the importance of aqueous-phase
672 heterogeneous reactions chemistry in the particle phase (Hennigan et al., 2009;
673 Volkamer et al., 2009). No significant correlation was observed between \log
674 $(K_{H,w}/K_{H,e})$ and NH_4^+ or NO_3^- concentrations. Although several studies have estimated
675 Henry’s law constants for a variety of polar organic compounds in pure water (e.g.,

676 ~~polyols and polyacids; Compernelle and Müller, 2014a, b), more work is warranted to~~
677 ~~decrease the estimation uncertainty and explain their increased partitioning toward~~
678 ~~aerosol liquid water explicitly. Therefore, the bulk WSOC and sulfate ion should play~~
679 ~~important roles during the condensation of gas-phase polyols, and further research is~~
680 ~~warranted to explicitly explain these effects.~~

681 **4 Implications and Conclusions and implications**

682 In this work, concentrations of gas- and particle-phase polyol tracers were
683 measured simultaneously in northern Nanjing. The temporal variations of individual
684 compounds were dominated by their particle-phase concentrations. ~~Because receptor-~~
685 ~~based models identify and quantify aerosol sources based on inter-sample variability,~~
686 Then gas-particle partitioning of polyol tracers should have little influence on source
687 apportionment ~~barely using based on~~ particle-phase data ~~in northern~~ Nanjing. ~~When it~~
688 ~~comes to other places (e.g., western US) where the concentration of aerosol liquid~~
689 ~~water is extremely low, the influence of gas-particle partitioning will still be a concern.~~
690 An improved agreement between measurement-based and predicted $K_{p,OM}$ of polyol
691 tracers was observed when the solubility in aerosol liquid water was considered,
692 indicating that the aqueous solution in aerosols is also an important absorbing phase.
693 ~~Similar to southeastern US, the ambient atmosphere in northern Nanjing also favored~~
694 ~~the condensation of polyol tracers, which was ascribed to the significant ALWC in~~
695 ~~these locations. The large gaps of $K_{H,c}^m$ versus $K_{H,w}^l$ measured versus predicted $K_{p,OM}$~~
696 ~~and K_H implied could be partly parameterized using the equation defining “salting-in”~~
697 ~~effects. According to existing studies, that the gas-particle partitioning of polyol~~
698 ~~tracers could not be depicted using equilibrium absorptive partitioning model or~~
699 ~~Henry’s law alone. In addition to the “salting-in” effect primarily due to the sulfate~~

700 ~~ions, other aerosol components like bulk WSOC~~reactive uptake, aqueous phase
701 reactions, and chemical similarity between partitioning species and the absorbing
702 phase might ~~also~~ be responsible for increasing the partitioning of polyol tracers into
703 the condensed phase. So, the results of this study have important implications on the
704 prediction of gas-particle partitioning of water-soluble organics, and further studies
705 are required to explain their enhanced aqueous-phase uptake mechanistically. Due to
706 the hygroscopic properties of highly oxidized organic aerosols, ~~the proposed scheme~~
707 ~~for gas-particle partitioning of polyol tracer~~this study also partly reveals the
708 discrepancy between modeled and observed SOA in previous studies. However,
709 several ~~pre~~-assumptions (e.g., ~~liquid-liquid phase separation~~LLPS) were made for ~~the~~
710 proposed gas-particle partitioning schemes in this work, more laboratory research is
711 needed to understand the mixing state of inorganic salts, organic components, and
712 aerosol liquid water in atmospheric particles.

714 *Data availability*

715 Data used in the writing of this paper is available at the Harvard Dataverse
716 (<https://doi.org/10.7910/DVN/U3IGQR>, Qin et al., 2021)

718 *Author contributions*

719 MX designed the research. CQ and YG performed the sampling and chemical analysis.
720 CQ, YM, and MX analyzed the data. CQ and MX wrote the paper with significant
721 contributions from YW, HL, and QW.

723 *Competing interests*

724 The authors declare that they have no conflict of interest.

725

726 *Acknowledgements*

727 This research was supported by the National Natural Science Foundation of China
728 (NSFC, 41701551). Y. W. was supported by the National Science Foundation
729 Atmospheric Chemistry Program.

730

731 **References**

- 732 [Ahrens, L., Shoeib, M., Harner, T., Lane, D. A., Guo, R., and Reiner, E. J.: Comparison of annular](#)
733 [diffusion denuder and high volume air samplers for measuring per- and polyfluoroalkyl](#)
734 [substances in the atmosphere, *Anal. Chem.*, 83, 9622-9628, 10.1021/ac202414w, 2011.](#)
- 735 Almeida, M. B., Alvarez, A. M., Miguel, E. M. D., and Hoyo, E. S. D.: Setchenow coefficients for
736 naphthols by distribution method, *Can. J. Chem.*, 61, 244-248, 10.1139/v83-043, 1983.
- 737 [Barsanti, K. C., and Pankow, J. F.: Thermodynamics of the formation of atmospheric organic](#)
738 [particulate matter by accretion reactions—Part 1: aldehydes and ketones, *Atmos. Environ.*, 38,](#)
739 [4371-4382, 10.1016/j.atmosenv.2004.03.035, 2004.](#)
- 740 Bauer, H., Claeys, M., Vermeylen, R., Schueller, E., Weinke, G., Berger, A., and Puxbaum, H.:
741 Arabitol and mannitol as tracers for the quantification of airborne fungal spores, *Atmos. Environ.*,
742 42, 588-593, <https://doi.org/10.1016/j.atmosenv.2007.10.013>, 2008.
- 743 Burshtein, N., Lang-Yona, N., and Rudich, Y.: Ergosterol, arabitol and mannitol as tracers for biogenic
744 aerosols in the eastern Mediterranean, *Atmos. Chem. Phys.*, 11, 829-839, 10.5194/acp-11-829-
745 2011, 2011.
- 746 [Carlton, A. G., and Turpin, B. J.: Particle partitioning potential of organic compounds is highest in the](#)
747 [Eastern US and driven by anthropogenic water, *Atmos. Chem. Phys.*, 13, 10203-10214,](#)
748 [10.5194/acp-13-10203-2013, 2013.](#)
- 749 [Chow, J. C., Watson, J. G., Chen, L. W. A., Rice, J., and Frank, N. H.: Quantification of PM_{2.5} organic](#)
750 [carbon sampling artifacts in US networks, *Atmos. Chem. Phys.*, 10, 5223-5239, 10.5194/acp-10-](#)
751 [5223-2010, 2010.](#)
- 752 Claeys, M., Graham, B., Vas, G., Wang, W., Vermeylen, R., Pashynska, V., Cafmeyer, J., Guyon, P.,
753 Andreae, M. O., Artaxo, P., and Maenhaut, W.: Formation of secondary organic aerosols through
754 photooxidation of isoprene, *Science*, 303, 1173-1176, 10.1126/science.1092805, 2004.
- 755 Compernelle, S., and Müller, J. F.: Henry's law constants of polyols, *Atmos. Chem. Phys.*, 14, 12815-
756 12837, 10.5194/acp-14-12815-2014, 2014a.
- 757 Compernelle, S., and Müller, J. F.: Henry's law constants of diacids and hydroxy polyacids:
758 recommended values, *Atmos. Chem. Phys.*, 14, 2699-2712, 10.5194/acp-14-2699-2014, 2014b.
- 759 [Cui, T., Zeng, Z., dos Santos, E. O., Zhang, Z., Chen, Y., Zhang, Y., Rose, C. A., Budisulistiorini, S. H.,](#)
760 [Collins, L. B., Bodnar, W. M., de Souza, R. A. F., Martin, S. T., Machado, C. M. D., Turpin, B. J.,](#)
761 [Gold, A., Ault, A. P., and Surratt, J. D.: Development of a hydrophilic interaction liquid](#)
762 [chromatography \(HILIC\) method for the chemical characterization of water-soluble isoprene](#)
763 [epoxydiol \(IEPOX\)-derived secondary organic aerosol, *Environmental Science: Environ. Sci.:*](#)
764 [Process. Impacts](#), 20, 1524-1536, 10.1039/C8EM00308D, 2018.
- 765 Fu, P., and Kawamura, K.: Diurnal variations of polar organic tracers in summer forest aerosols: A case
766 study of a Quercus and Picea mixed forest in Hokkaido, Japan, *Geochem. J.*, 45, 297-308,
767 10.2343/geochemj.1.0123, 2011.
- 768 Galloway, M. M., Chhabra, P. S., Chan, A. W. H., Surratt, J. D., Flagan, R. C., Seinfeld, J. H., and
769 Keutsch, F. N.: Glyoxal uptake on ammonium sulphate seed aerosol: reaction products and
770 reversibility of uptake under dark and irradiated conditions, *Atmos. Chem. Phys.*, 9, 3331-3345,

771 10.5194/acp-9-3331-2009, 2009.

772 Hallar, A. G., Lowenthal, D. H., Clegg, S. L., Samburova, V., Taylor, N., Mazzoleni, L. R., Zielinska,
773 B. K., Kristensen, T. B., Chirokova, G., McCubbin, I. B., Dodson, C., and Collins, D.: Chemical
774 and hygroscopic properties of aerosol organics at Storm Peak Laboratory, *J. Geophys. Res.*
775 *Atmos.*, 118, 4767-4779, <https://doi.org/10.1002/jgrd.50373>, 2013.

776 Heald, C. L., Jacob, D. J., Park, R. J., Russell, L. M., Huebert, B. J., Seinfeld, J. H., Liao, H., and
777 Weber, R. J.: A large organic aerosol source in the free troposphere missing from current models,
778 *Geophys. Res. Lett.*, 32, L18809, <https://doi.org/10.1029/2005GL023831>, 2005.

779 Healy, R. M., Wenger, J. C., Metzger, A., Duplissy, J., Kalberer, M., and Dommen, J.: Gas/particle
780 partitioning of carbonyls in the photooxidation of isoprene and 1,3,5-trimethylbenzene, *Atmos.*
781 *Chem. Phys.*, 8, 3215-3230, 10.5194/acp-8-3215-2008, 2008.

782 Hennigan, C. J., Bergin, M. H., Russell, A. G., Nenes, A., and Weber, R. J.: Gas/particle partitioning of
783 water-soluble organic aerosol in Atlanta, *Atmos. Chem. Phys.*, 9, 3613-3628, 10.5194/acp-9-
784 3613-2009, 2009.

785 Hennigan, C. J., Sullivan, A. P., Collett Jr, J. L., and Robinson, A. L.: Levoglucosan stability in
786 biomass burning particles exposed to hydroxyl radicals, *Geophys. Res. Lett.*, 37, L09806,
787 <https://doi.org/10.1029/2010GL043088>, 2010.

788 [Hilal, S. H., Ayyampalayam, S. N., and Carreira, L. A.: Air-liquid partition coefficient for a diverse set](#)
789 [of organic compounds: Henry's Law constant in water and hexadecane, *Environ. Sci. Technol.*,](#)
790 [42, 9231-9236, 10.1021/es8005783, 2008.](#)

791 Hodzic, A., Jimenez, J. L., Madronich, S., Canagaratna, M. R., DeCarlo, P. F., Kleinman, L., and Fast,
792 J.: Modeling organic aerosols in a megacity: potential contribution of semi-volatile and
793 intermediate volatility primary organic compounds to secondary organic aerosol formation,
794 *Atmos. Chem. Phys.*, 10, 5491-5514, 10.5194/acp-10-5491-2010, 2010.

795 Hoffmann, D., Tilgner, A., Iinuma, Y., and Herrmann, H.: Atmospheric stability of levoglucosan: A
796 detailed laboratory and modeling study, *Environ. Sci. Technol.*, 44, 694-699, 10.1021/es902476f,
797 2010.

798 Hu, D., Bian, Q., Lau, A. K. H., and Yu, J. Z.: Source apportioning of primary and secondary organic
799 carbon in summer PM_{2.5} in Hong Kong using positive matrix factorization of secondary and
800 primary organic tracer data, *J. Geophys. Res. Atmos.*, 115, D16204,
801 <https://doi.org/10.1029/2009JD012498>, 2010.

802 Ip, H. S. S., Huang, X. H. H., and Yu, J. Z.: Effective Henry's law constants of glyoxal, glyoxylic acid,
803 and glycolic acid, *Geophys. Res. Lett.*, 36, L01802, <https://doi.org/10.1029/2008GL036212>, 2009.

804 Isaacman-VanWertz, G., Yee, L. D., Kreisberg, N. M., Wernis, R., Moss, J. A., Hering, S. V., de Sá, S.
805 S., Martin, S. T., Alexander, M. L., Palm, B. B., Hu, W., Campuzano-Jost, P., Day, D. A.,
806 Jimenez, J. L., Riva, M., Surratt, J. D., Viegas, J., Manzi, A., Edgerton, E., Baumann, K., Souza,
807 R., Artaxo, P., and Goldstein, A. H.: Ambient gas-particle partitioning of tracers for biogenic
808 oxidation, *Environ. Sci. Technol.*, 50, 9952-9962, 10.1021/acs.est.6b01674, 2016.

809 Jang, M., Czoschke, N. M., Lee, S., and Kamens, R. M.: Heterogeneous atmospheric aerosol
810 production by acid-catalyzed particle-phase reactions, *Science*, 298, 814,
811 10.1126/science.1075798, 2002.

812 Kampf, C. J., Waxman, E. M., Slowik, J. G., Dommen, J., Pfaffenberger, L., Praplan, A. P., Prévôt, A.
813 S. H., Baltensperger, U., Hoffmann, T., and Volkamer, R.: Effective Henry's law partitioning and
814 the salting constant of glyoxal in aerosols containing sulfate, *Environ. Sci. Technol.*, 47, 4236-
815 4244, 10.1021/es400083d, 2013.

816 Kondo, Y., Miyazaki, Y., Takegawa, N., Miyakawa, T., Weber, R. J., Jimenez, J. L., Zhang, Q., and
817 Worsnop, D. R.: Oxygenated and water-soluble organic aerosols in Tokyo, *J. Geophys. Res.*
818 *Atmos.*, 112, D01203, 10.1029/2006jd007056, 2007.

819 Kroll, J. H., Ng, N. L., Murphy, S. M., Varutbangkul, V., Flagan, R. C., and Seinfeld, J. H.: Chamber
820 studies of secondary organic aerosol growth by reactive uptake of simple carbonyl compounds, *J.*
821 *Geophys. Res. Atmos.*, 110, D23207, <https://doi.org/10.1029/2005JD006004>, 2005.

822 [Lin, Y.-H., Budisulistiorini, S. H., Chu, K., Siejack, R. A., Zhang, H., Riva, M., Zhang, Z., Gold, A.,](#)
823 [Kautzman, K. E., and Surratt, J. D.: Light-absorbing oligomer formation in secondary organic](#)
824 [aerosol from reactive uptake of isoprene epoxydiols, *Environ. Sci. Technol.*, 48, 12012-12021,](#)
825 [10.1021/es503142b, 2014.](#)

826 Lin, Y. H., Knipping, E. M., Edgerton, E. S., Shaw, S. L., and Surratt, J. D.: Investigating the
827 influences of SO₂ and NH₃ levels on isoprene-derived secondary organic aerosol formation using
828 conditional sampling approaches, *Atmos. Chem. Phys.*, 13, 8457-8470, 10.5194/acp-13-8457-

829 2013, 2013.

830 Liu, J., Li, X., Li, D., Xu, R., Gao, Y., Chen, S., Liu, Y., Zhao, G., Wang, H., Wang, H., Lou, S., Chen,
831 M., Hu, J., Lu, K., Wu, Z., Hu, M., Zeng, L., and Zhang, Y.: Observations of glyoxal and
832 methylglyoxal in a suburban area of the Yangtze River Delta, China, *Atmos. Environ.*, 238,
833 117727, <https://doi.org/10.1016/j.atmosenv.2020.117727>, 2020.

834 Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.: Gas/particle
835 partitioning and secondary organic aerosol yields, *Environ. Sci. Technol.*, 30, 2580-2585,
836 10.1021/es950943+, 1996.

837 Pankow, J. F.: An absorption model of the gas/aerosol partitioning involved in the formation of
838 secondary organic aerosol, *Atmos. Environ.*, 28, 189-193, 10.1016/1352-2310(94)90094-9, 1994a.

839 Pankow, J. F.: An absorption model of gas/particle partitioning of organic compounds in the
840 atmosphere, *Atmos. Environ.*, 28, 185-188, 10.1016/1352-2310(94)90093-0, 1994b.

841 Perraud, V., Bruns, E. A., Ezell, M. J., Johnson, S. N., Yu, Y., Alexander, M. L., Zelenyuk, A., Imre,
842 D., Chang, W. L., Dabdub, D., Pankow, J. F., and Finlayson-Pitts, B. J.: Nonequilibrium
843 atmospheric secondary organic aerosol formation and growth, *Proc. Natl. Acad. Sci. U.S.A.*, 109,
844 2836-2841, 10.1073/pnas.1119909109, 2012.

845 Peters, A. J., Lane, D. A., Gundel, L. A., Northcott, G. L., and Jones, K. C.: A comparison of high
846 volume and diffusion denuder samplers for measuring semivolatile organic compounds in the
847 atmosphere, *Environ. Sci. Technol.*, 34, 5001-5006, 10.1021/es000056t, 2000.

848 ~~Pye, H. O. T., Zuend, A., Fry, J. L., Isaacman-VanWertz, G., Capps, S. L., Appel, K. W., Foroutan, H.,
849 Xu, L., Ng, N. L., and Goldstein, A. H.: Coupling of organic and inorganic aerosol systems and
850 the effect on gas-particle partitioning in the southeastern US, *Atmos. Chem. Phys.*, 18, 357-370,
851 10.5194/acp-18-357-2018, 2018.~~

852 Qin, C., Gou, Y., Wang, Y., Mao, Y., Liao, H., Wang, Q. g., and Xie, M.: Replication Data for: Gas-
853 particle partitioning of polyol tracers in the western Yangtze River Delta, China, in, V1 ed.,
854 Harvard Dataverse, <https://doi.org/10.7910/DVN/U3IGQR>, 2021.

855 Sadeghi, R., and Jahani, F.: Salting-in and salting-out of water-soluble polymers in aqueous salt
856 solutions, *J. Phys. Chem. B*, 116, 5234-5241, 10.1021/jp300665b, 2012.

857 Saxena, P., and Hildemann, L.: Water-soluble organics in atmospheric particles: A critical review of
858 the literature and application of thermodynamics to identify candidate compounds, *J. Atmos.
859 Chem.*, 24, 57-109, 10.1007/bf00053823, 1996.

860 ~~Seinfeld, J. H., and Pandis, S. N.: Atmospheric chemistry and physics: from air pollution to climate
861 change, John Wiley & Sons, 2016.~~

862 Setschenow, J.: Über die Konstitution der Salzlösungen auf Grund ihres Verhaltens zu Kohlensäure, *Z.
863 Phys. Chem.*, 4U, 117-125, <https://doi.org/10.1515/zpch-1889-0409>, 1889.

864 Shen, H., Chen, Z., Li, H., Qian, X., Qin, X., and Shi, W.: Gas-particle partitioning of carbonyl
865 compounds in the ambient atmosphere, *Environ. Sci. Technol.*, 52, 10997-11006,
866 10.1021/acs.est.8b01882, 2018.

867 Simcik, M. F., Franz, T. P., Zhang, H., and Eisenreich, S. J.: Gas-particle partitioning of PCBs and
868 PAHs in the Chicago urban and adjacent coastal atmosphere: States of equilibrium, *Environ. Sci.
869 Technol.*, 32, 251-257, 10.1021/es970557n, 1998.

870 Simoneit, B. R. T., Schauer, J. J., Nolte, C. G., Oros, D. R., Elias, V. O., Fraser, M. P., Rogge, W. F.,
871 and Cass, G. R.: Levoglucosan, a tracer for cellulose in biomass burning and atmospheric
872 particles, *Atmos. Environ.*, 33, 173-182, [http://dx.doi.org/10.1016/S1352-2310\(98\)00145-9](http://dx.doi.org/10.1016/S1352-2310(98)00145-9), 1999.

873 Simoneit, B. R. T., Elias, V. O., Kobayashi, M., Kawamura, K., Rushdi, A. I., Medeiros, P. M., Rogge,
874 W. F., and Didyk, B. M.: Sugars dominant water-soluble organic compounds in soils and
875 characterization as tracers in atmospheric particulate matter, *Environ. Sci. Technol.*, 38, 5939-
876 5949, 10.1021/es0403099, 2004.

877 ~~Subramanian, R., Khlystov, A. Y., Cabada, J. C., and Robinson, A. L.: Positive and negative artifacts in
878 particulate organic carbon measurements with denuded and undenuded sampler configurations
879 special issue of aerosol science and technology on findings from the fine particulate matter
880 supersites program, *Aerosol Sci. Tech.*, 38, 27-48, 10.1080/02786820390229354, 2004.~~

881 Surratt, J. D., Murphy, S. M., Kroll, J. H., Ng, N. L., Hildebrandt, L., Sorooshian, A., Szmigielski, R.,
882 Vermeylen, R., Maenhaut, W., Claeys, M., Flagan, R. C., and Seinfeld, J. H.: Chemical
883 composition of secondary organic aerosol formed from the photooxidation of isoprene, *J. Phys.
884 Chem. A*, 110, 9665-9690, 10.1021/jp061734m, 2006.

885 Surratt, J. D., Chan, A. W. H., Eddingsaas, N. C., Chan, M., Loza, C. L., Kwan, A. J., Hersey, S. P.,
886 Flagan, R. C., Wennberg, P. O., and Seinfeld, J. H.: Reactive intermediates revealed in secondary

887 organic aerosol formation from isoprene, *Proc. Natl. Acad. Sci. U.S.A.*, 107, 6640-6645,
888 10.1073/pnas.0911114107, 2010.

889 Taylor, N. F., Collins, D. R., Lowenthal, D. H., McCubbin, I. B., Hallar, A. G., Samburova, V.,
890 Zielinska, B., Kumar, N., and Mazzoleni, L. R.: Hygroscopic growth of water soluble organic
891 carbon isolated from atmospheric aerosol collected at US national parks and Storm Peak
892 Laboratory, *Atmos. Chem. Phys.*, 17, 2555-2571, 10.5194/acp-17-2555-2017, 2017.

893 Tsyro, S. G.: To what extent can aerosol water explain the discrepancy between model calculated and
894 gravimetric PM10 and PM2.5?, *Atmos. Chem. Phys.*, 5, 515-532, 10.5194/acp-5-515-2005, 2005.

895 Turpin, B. J., and Lim, H.-J.: Species contributions to PM2.5 mass concentrations: Revisiting common
896 assumptions for estimating organic mass, *Aerosol Sci. Technol.*, 35, 602-610,
897 10.1080/02786820119445, 2001.

898 US EPA: Estimation Program Interface (EPI) Suite. Version v4.11, 2012. United States Environmental
899 Protection Agency, Washington, DC, USA. [https://www.epa.gov/tsca-screening-tools/download-](https://www.epa.gov/tsca-screening-tools/download-epi-suite-estimation-program-interface-v411)
900 [epi-suite-estimation-program-interface-v411](https://www.epa.gov/tsca-screening-tools/download-epi-suite-estimation-program-interface-v411)

901 Volkamer, R., Jimenez, J. L., San Martini, F., Dzepina, K., Zhang, Q., Salcedo, D., Molina, L. T.,
902 Worsnop, D. R., and Molina, M. J.: Secondary organic aerosol formation from anthropogenic air
903 pollution: Rapid and higher than expected, *Geophys. Res. Lett.*, 33, L17811,
904 <https://doi.org/10.1029/2006GL026899>, 2006.

905 Volkamer, R., Ziemann, P. J., and Molina, M. J.: Secondary Organic Aerosol Formation from
906 Acetylene (C2H2): seed effect on SOA yields due to organic photochemistry in the aerosol
907 aqueous phase, *Atmos. Chem. Phys.*, 9, 1907-1928, 10.5194/acp-9-1907-2009, 2009.

908 Wang, W., Kourtchev, I., Graham, B., Cafmeyer, J., Maenhaut, W., and Claeys, M.: Characterization
909 of oxygenated derivatives of isoprene related to 2-methyltetrols in Amazonian aerosols using
910 trimethylsilylation and gas chromatography/ion trap mass spectrometry, *Rapid Commun. Mass*
911 *Spectrom.*, 19, 1343-1351, <https://doi.org/10.1002/rcm.1940>, 2005.

912 Waxman, E. M., Elm, J., Kurtén, T., Mikkelsen, K. V., Ziemann, P. J., and Volkamer, R.: Glyoxal and
913 methylglyoxal Setschenow salting constants in sulfate, nitrate, and chloride solutions:
914 Measurements and Gibbs energies, *Environ. Sci. Technol.*, 49, 11500-11508,
915 10.1021/acs.est.5b02782, 2015.

916 Williams, B. J., Goldstein, A. H., Kreisberg, N. M., and Hering, S. V.: In situ measurements of
917 gas/particle-phase transitions for atmospheric semivolatile organic compounds, *Proc. Natl. Acad.*
918 *Sci. U.S.A.*, 107, 6676-6681, 10.1073/pnas.0911858107, 2010.

919 Xie, M., Barsanti, K. C., Hannigan, M. P., Dutton, S. J., and Vedal, S.: Positive matrix factorization of
920 PM2.5 - eliminating the effects of gas/particle partitioning of semivolatile organic compounds,
921 *Atmos. Chem. Phys.*, 13, 7381-7393, 10.5194/acp-13-7381-2013, 2013.

922 Xie, M., Hannigan, M. P., and Barsanti, K. C.: Gas/particle partitioning of n-alkanes, PAHs and
923 oxygenated PAHs in urban Denver, *Atmos. Environ.*, 95, 355-362,
924 <http://dx.doi.org/10.1016/j.atmosenv.2014.06.056>, 2014a.

925 Xie, M., Hannigan, M. P., and Barsanti, K. C.: Gas/particle partitioning of 2-methyltetrols and
926 levoglucosan at an urban site in Denver, *Environ. Sci. Technol.*, 48, 2835-2842,
927 10.1021/es405356n, 2014b.

928 Xie, M., Hannigan, M. P., and Barsanti, K. C.: Impact of gas/particle partitioning of semivolatile
929 organic compounds on source apportionment with positive matrix factorization, *Environ. Sci.*
930 *Technol.*, 48, 9053-9060, 10.1021/es5022262, 2014c.

931 Yang, L., Shang, Y., Hannigan, M. P., Zhu, R., Wang, Q. g., Qin, C., and Xie, M.: Collocated
932 speciation of PM2.5 using tandem quartz filters in northern nanjing, China: Sampling artifacts and
933 measurement uncertainty, *Atmos. Environ.*, 246, 118066,
934 <https://doi.org/10.1016/j.atmosenv.2020.118066>, 2021.

935 Yatavelli, R. L. N., Stark, H., Thompson, S. L., Kimmel, J. R., Cubison, M. J., Day, D. A.,
936 Campuzano-Jost, P., Palm, B. B., Hodzic, A., Thornton, J. A., Jayne, J. T., Worsnop, D. R., and
937 Jimenez, J. L.: Semicontinuous measurements of gas-particle partitioning of organic acids in a
938 ponderosa pine forest using a MOVI-HRToF-CIMS, *Atmos. Chem. Phys.*, 14, 1527-1546,
939 10.5194/acp-14-1527-2014, 2014.

940 Zhang, Y., Sheesley, R. J., Schauer, J. J., Lewandowski, M., Jaoui, M., Offenber, J. H., Kleindienst, T.
941 E., and Edney, E. O.: Source apportionment of primary and secondary organic aerosols using
942 positive matrix factorization (PMF) of molecular markers, *Atmos. Environ.*, 43, 5567-5574,
943 <https://doi.org/10.1016/j.atmosenv.2009.02.047>, 2009.

944 Zhao, Y., Kreisberg, N. M., Worton, D. R., Isaacman, G., Weber, R. J., Liu, S., Day, D. A., Russell, L.

945 M., Markovic, M. Z., VandenBoer, T. C., Murphy, J. G., Hering, S. V., and Goldstein, A. H.:
946 Insights into secondary organic aerosol formation mechanisms from measured gas/particle
947 partitioning of specific organic tracer compounds, *Environ. Sci. Technol.*, 47, 3781-3787,
948 10.1021/es304587x, 2013.
949 [Zuend, A., and Seinfeld, J. H.: Modeling the gas-particle partitioning of secondary organic aerosol: the
950 importance of liquid-liquid phase separation, *Atmos. Chem. Phys.*, 12, 3857-3882, 10.5194/acp-
951 12-3857-2012, 2012.](#)
952

Table 1. Comparisons of measurement-based $\log K_{p,OM}$ ($m^3 \mu g^{-1}$) at three proposed cases and predicted values.

Species	No. of obs.	$\log K_{p,OM}^m$ ^a			$\log K_{p,OM}^t$ ^b			
		Case 1	Case 2	Case 3	EPI	EVAPORATION	SPARC	SIMPOL
Isoprene SOA tracers								
C5-alkene triol 1	53	0.33 ± 0.71	-0.79 ± 0.86	-0.82 ± 0.85	-3.09	-2.84	-1.19	-2.88
C5-alkene triol 2	63	0.15 ± 0.55	-1.02 ± 0.74	-1.05 ± 0.73	-3.62	-3.67	-4.14	-2.85
C5-alkene triol 3	83	0.35 ± 0.68	-0.83 ± 0.86	-0.86 ± 0.85	-2.90	-2.65	-1.00	-2.69
2-Methylthreitol	101	-0.12 ± 0.48	-2.09 ± 0.71	-2.09 ± 0.70	-1.87	-1.30	-1.18	-0.47
2-Methylerythritol	95	-0.011 ± 0.58	-1.96 ± 0.71	-1.96 ± 0.71	-1.90	-1.34	-1.22	-0.50
Biomass burning tracer								
Levoglucozan	65	2.23 ± 0.72	0.63 ± 0.90	0.62 ± 0.90	-0.04	-0.81	1.04	-0.76
Sugars and sugar alcohols								
Meso-erythritol	31	0.87 ± 0.53	-1.43 ± 0.60	-1.43 ± 0.60	-0.65	-1.21	-0.45	
Fructose	85	0.65 ± 0.73	-1.20 ± 0.83	-1.20 ± 0.89	1.17	2.76	6.94	
Mannose	74	0.62 ± 0.71	-2.12 ± 0.95	-2.12 ± 0.95	1.28	2.13	4.77	
Glucose	88	0.42 ± 0.67	-2.77 ± 0.93	-2.77 ± 0.93	0.34	3.75	7.32	
Xylitol	22	0.24 ± 0.54	-2.61 ± 0.72	-2.61 ± 0.72	3.37	2.34	3.57	
Arabitel	30	1.46 ± 0.89	-1.35 ± 1.24	-1.35 ± 1.24	3.25	1.67	2.90	
Manitol	65	1.08 ± 0.63	-2.24 ± 0.95	-2.24 ± 0.95	2.33	4.16	6.68	

^a Average ± standard deviation; ^b temperature range: -4~36 °C.

Table 1. Statistics for measured and predicted $\log K_{p,OM}$ of individual polyol tracers.

Species	No. of obs.	$\log K_{p,OM}^m$			$\log K_{p,OM}^t$		
		Median	Average	Range	Median	Average	Range
Isoprene SOA tracers							
C5-alkene triol 1	53	0.47	0.33 ± 0.71	-1.30—2.61	-3.23	-3.09 ± 0.44	-3.70—1.63
C5-alkene triol 2	63	0.19	0.15 ± 0.55	-1.02—2.26	-3.70	-3.62 ± 0.37	-4.24—-2.67
C5-alkene triol 3	83	0.38	0.35 ± 0.68	-1.86—2.25	-3.01	-2.90 ± 0.48	-3.70—1.63
2-Methylthreitol	101	-0.13	-0.12 ± 0.48	-1.15—0.92	-1.90	-1.87 ± 0.52	-2.81—-0.62
2-Methylerythritol	95	-0.089	-0.011 ± 0.58	-1.09—2.06	-1.91	-1.90 ± 0.51	-2.81—-0.62
Biomass burning tracer							
Levoglucozan	65	2.34	2.23 ± 0.72	-0.11—3.36	-0.12	-0.038 ± 0.59	-1.00—1.29
Sugars and sugar alcohols							
Meso-erythritol	31	0.84	0.87 ± 0.53	-0.47—1.81	-0.80	-0.65 ± 0.48	-1.35—0.69
Fructose	85	0.55	0.65 ± 0.73	-0.96—3.01	1.14	1.17 ± 0.62	0.015—2.72
Mannose	74	0.57	0.62 ± 0.71	-0.94—2.81	1.23	1.28 ± 0.66	0.18—2.81
Glucose	88	0.41	0.42 ± 0.67	-1.08—1.92	0.31	0.34 ± 0.65	-0.75—1.92
Xylitol	22	0.35	0.24 ± 0.54	-1.23—0.97	3.43	3.37 ± 0.57	2.11—4.39
Arabitel	30	1.40	1.46 ± 0.89	-0.19—4.20	3.15	3.25 ± 0.77	2.05—4.81
Manitol	65	1.06	1.08 ± 0.63	-0.35—2.53	2.31	2.33 ± 0.70	1.15—3.98

Table 2. Comparisons of measurement-based $\log K_{H,e}$ ($\text{mol m}^{-3} \text{atm}^{-1}$) and predicted $\log K_{H,w}$ of individual polyol tracers.

Species	No. of obs.	Log $K_{H,e}^m$ (Cases 2) ^a			Log $K_{H,w}^b$	
		Median	Average	Range	EPI	SPARC
<u>Isoprene SOA tracers</u>						
<u>C5-alkene triol 1</u>	<u>53</u>	<u>14.0</u>	<u>13.9 ± 0.86</u>	<u>11.5 – 16.4</u>	<u>7.22</u>	<u>11.7</u>
<u>C5-alkene triol 2</u>	<u>63</u>	<u>13.7</u>	<u>13.6 ± 0.73</u>	<u>11.2 – 16.1</u>	<u>7.34</u>	<u>7.66</u>
<u>C5-alkene triol 3</u>	<u>83</u>	<u>13.9</u>	<u>13.8 ± 0.85</u>	<u>10.6 – 16.1</u>	<u>7.43</u>	<u>11.9</u>
<u>2-Methylthreitol</u>	<u>101</u>	<u>13.4</u>	<u>13.3 ± 0.70</u>	<u>10.9 – 14.8</u>	<u>10.0</u>	<u>14.1</u>
<u>2-Methylerythritol</u>	<u>95</u>	<u>13.5</u>	<u>13.5 ± 0.71</u>	<u>11.6 – 15.6</u>	<u>9.95</u>	<u>14.1</u>
<u>Biomass burning tracer</u>						
<u>Levoglucofan</u>	<u>65</u>	<u>15.7</u>	<u>15.7 ± 0.90</u>	<u>13.2 – 17.3</u>	<u>13.4</u>	<u>16.1</u>
<u>Sugars and sugar alcohols</u>						
<u>Meso-erythritol</u>	<u>31</u>	<u>14.5</u>	<u>14.4 ± 0.60</u>	<u>12.8 – 15.6</u>	<u>9.65</u>	<u>13.8</u>
<u>Fructose</u>	<u>85</u>	<u>14.2</u>	<u>14.1 ± 0.89</u>	<u>11.9 – 16.5</u>	<u>14.7</u>	<u>19.9</u>
<u>Mannose</u>	<u>74</u>	<u>14.0</u>	<u>14.1 ± 0.94</u>	<u>12.1 – 16.8</u>	<u>10.9</u>	<u>18.8</u>
<u>Glucose</u>	<u>88</u>	<u>13.9</u>	<u>13.9 ± 0.93</u>	<u>11.3 – 16.3</u>	<u>14.7</u>	<u>20.9</u>
<u>Xylitol</u>	<u>22</u>	<u>13.8</u>	<u>13.7 ± 0.72</u>	<u>12.6 – 15.0</u>	<u>12.1</u>	<u>18.1</u>
<u>Arabitol</u>	<u>30</u>	<u>15.1</u>	<u>15.0 ± 1.23</u>	<u>13.0 – 18.2</u>	<u>11.3</u>	<u>17.4</u>
<u>Mannitol</u>	<u>65</u>	<u>14.6</u>	<u>14.5 ± 0.94</u>	<u>12.1 – 16.4</u>	<u>12.9</u>	<u>20.8</u>

^a Log $K_{H,e}^m$ values of Case 3 had ignorable difference, and were not exhibited separately; ^b temperature range: -4~36 °C

Table 2. Statistics for $\log K_{H,e}$ and $\log K_{H,w}$ of individual polyol tracers.

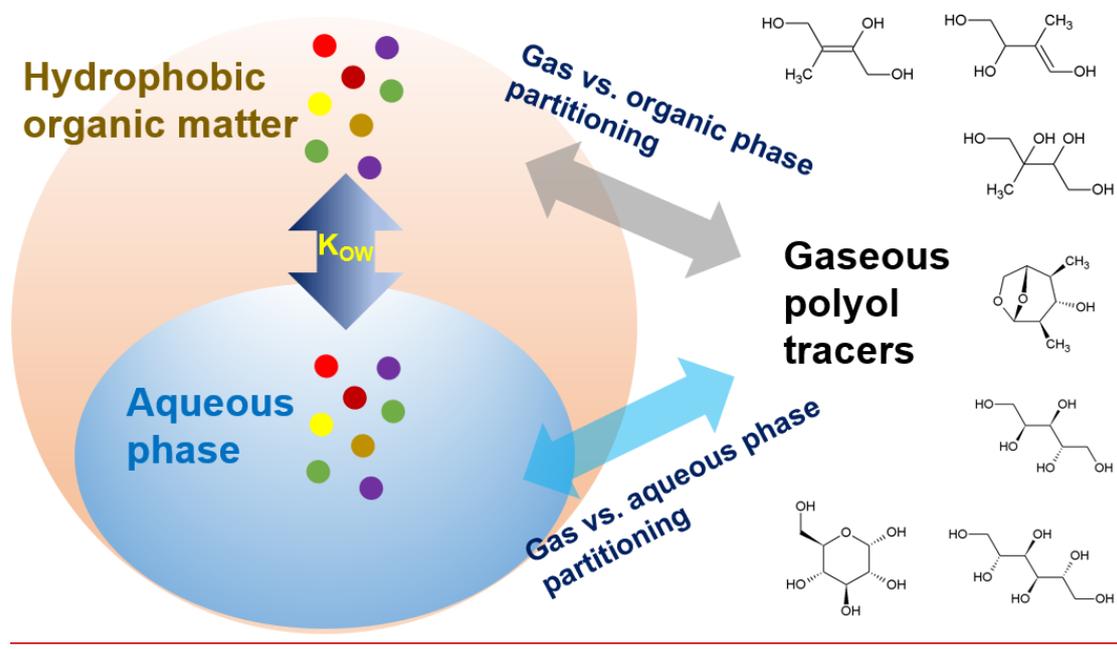
species	No. of obs.	Log $K_{H,e}$			Log $K_{H,w}$		
		Median	Average	Range	Median	Average	Range

<i>Isoprene-SOA tracers</i>							
C5-alkene-triol-1	53	14.0	13.9 ± 0.86	11.5—16.4	7.06	7.22 ± 0.50	6.53—8.87
C5-alkene-triol-2	63	13.7	13.6 ± 0.73	11.2—16.1	7.24	7.34 ± 0.45	6.60—8.49
C5-alkene-triol-3	83	13.9	13.8 ± 0.85	10.6—16.1	7.31	7.43 ± 0.55	6.53—8.87
2-Methylthreitol	101	13.4	13.3 ± 0.70	10.9—14.8	9.96	10.0 ± 0.80	8.55—11.9
2-Methylethritol	95	13.5	13.5 ± 0.71	11.6—15.6	9.93	9.95 ± 0.79	8.55—11.9
<i>Biomass-burning tracer</i>							
Levoglucozan	65	15.7	15.7 ± 0.90	13.2—17.3	13.3	13.4 ± 0.56	12.4—14.6
<i>Sugars and sugar alcohols</i>							
Meso-erythritol	31	14.5	14.4 ± 0.60	12.8—15.6	9.44	9.65 ± 0.66	8.68—11.5
Fructose	85	14.2	14.1 ± 0.89	11.9—16.5	14.6	14.7 ± 0.84	13.1—16.8
Mannose	74	14.0	14.1 ± 0.94	12.1—16.8	10.9	10.9 ± 0.88	9.46—13.0
Glucose	88	13.9	13.9 ± 0.93	11.3—16.3	14.6	14.7 ± 0.85	13.2—16.8
Xylitol	22	13.8	13.7 ± 0.72	12.6—15.0	12.1	12.1 ± 0.62	10.7—13.2
Arabitol	30	15.1	15.0 ± 1.23	13.0—18.2	11.2	11.4 ± 0.83	10.0—13.0
Mannitol	65	14.6	14.5 ± 0.94	12.1—16.4	12.9	12.9 ± 1.15	11.0—15.6

953

Figure 1

954



955

956

957

Figure 1. Proposed scheme for gas-particle partitioning of polyol tracers.

958

Figure 12

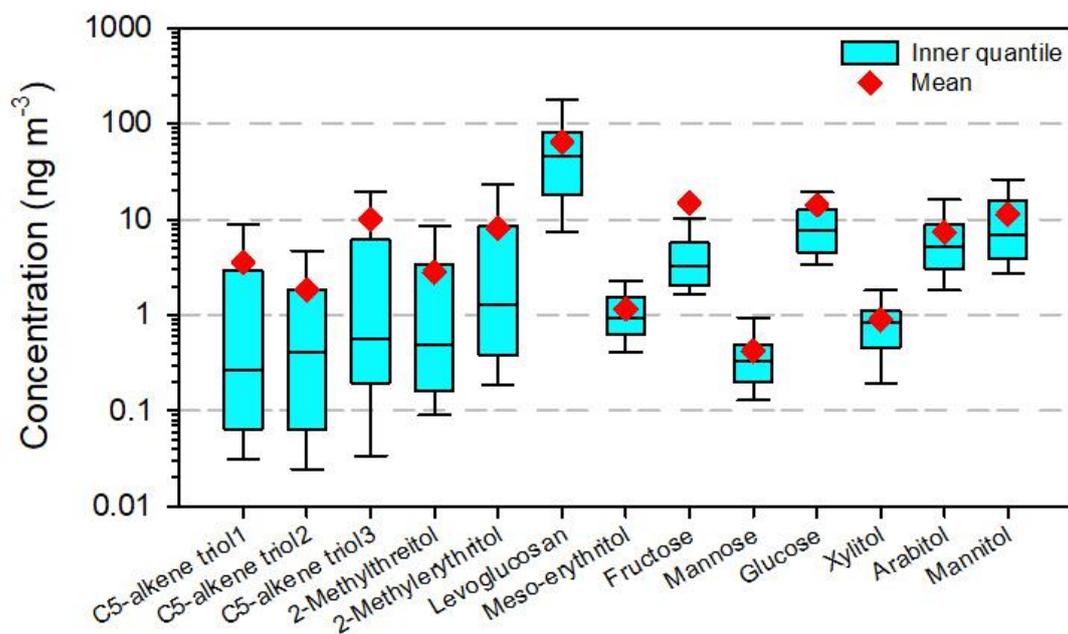
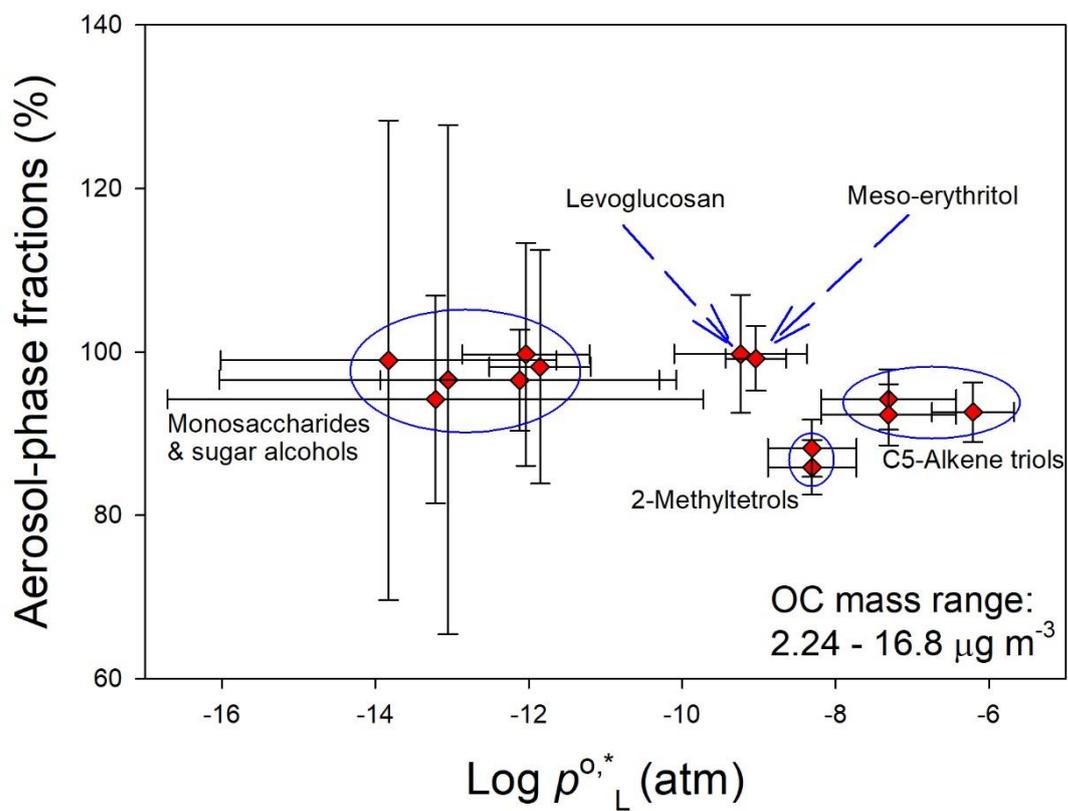


Figure 12. Total concentrations of individual polyols (Q_f + Q_b + PUF) in the ambient atmosphere of northern Nanjing. The boxes depict the median (dark line), inner quantile range (box), 10th and 90th percentiles (whiskers), and the mean (red diamond).

Figure 23



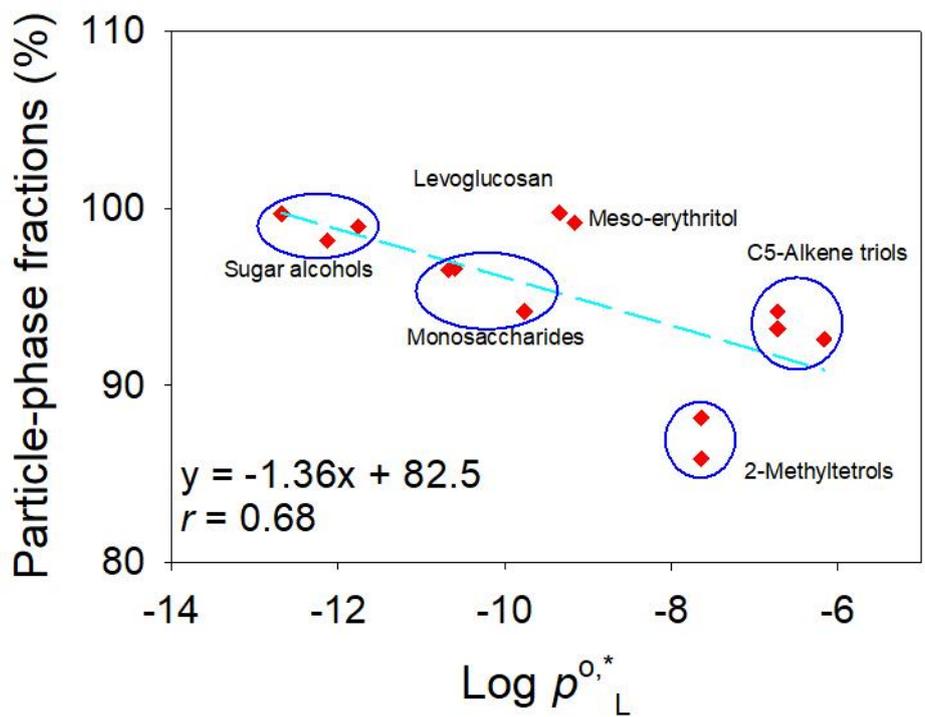
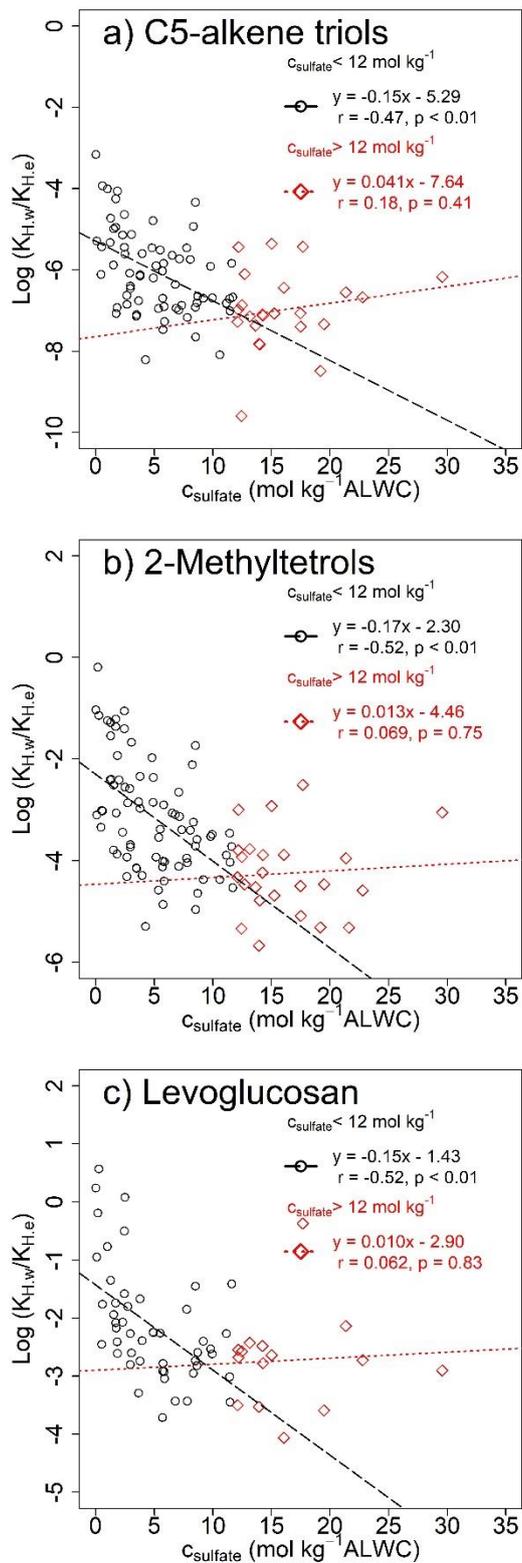


Figure 23. Average particle-phase fractions and $\log p^{0,*}_L$ of individual polyol tracers. Whiskers represent uncertainties of $F\%$ and one standard deviation of $\log p^{0,*}_L$ derived from different estimation tools.

~~Linear relationship between average aerosol-phase fractions and $\log p^{0,*}_L$ of polyol tracers.~~

Figure 34



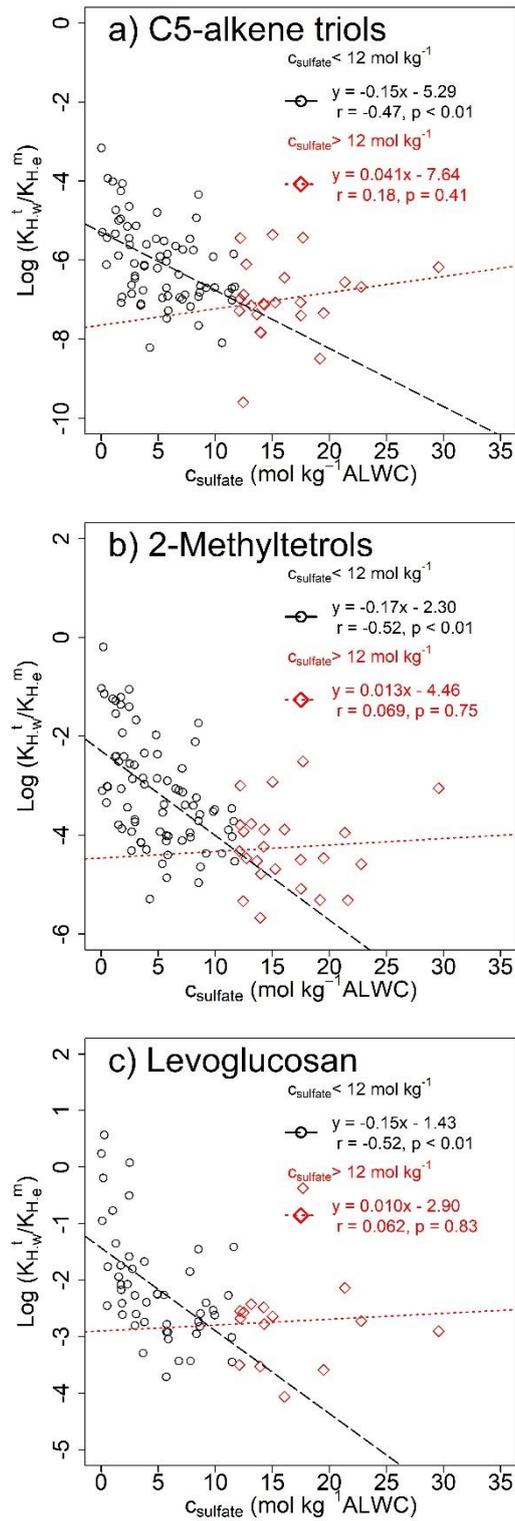


Figure 34. Modified Setschenow plots of $\log(K_{H,w}^1/K_{H,e}^m)$ vs. versus C_{sulfate} for (a) C5-alkene triols, (b) 2-methyltetrols, and (c) levoglucosan.