

# Isotopic evidence for dominant secondary production of HONO in near-ground wildfire plumes

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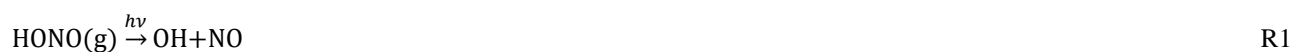
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**Abstract.** Nitrous acid (HONO) is an important precursor to hydroxyl radical (OH) that determines atmospheric oxidative  
15 capacity and thus impacts climate and air quality. Wildfire is not only a major direct source of HONO, it also results in highly  
polluted conditions that favour heterogeneous formation of HONO from nitrogen oxides ( $\text{NO}_x = \text{NO} + \text{NO}_2$ ) and nitrate on  
both ground and particle surfaces. However, these processes remain poorly constrained. To quantitatively constrain the HONO  
budget under various fire/smoke conditions, we combine a unique dataset of field concentrations and isotopic ratios ( $^{15}\text{N}/^{14}\text{N}$   
and  $^{18}\text{O}/^{16}\text{O}$ ) of  $\text{NO}_x$  and HONO, with an isotopic box model. Here we report the first isotopic evidence of secondary HONO  
20 production in near-ground wildfire plumes (over a sample integration time of hours), and the subsequent quantification of the  
relative importance of each pathway to total HONO production. Most importantly, our results reveal that nitrate photolysis  
plays a minor role (<5%) in HONO formation in daytime aged smoke, while  $\text{NO}_2$ -to-HONO heterogeneous conversion  
contributes 85-95% to total HONO production, followed by  $\text{OH}+\text{NO}$  (5-15%). At nighttime, heterogeneous reduction of  $\text{NO}_2$   
catalysed by redox active species (e.g., iron oxide and/or quinone) is essential ( $\geq 75\%$ ) for HONO production in addition to  
25 surface  $\text{NO}_2$  hydrolysis. Additionally, the  $^{18}\text{O}/^{16}\text{O}$  of HONO is used for the first time to constrain the  $\text{NO}$ -to- $\text{NO}_2$  oxidation  
branching ratio between ozone and peroxy radicals. Our approach provides a new and critical way to mechanistically constrain  
atmospheric chemistry/air quality models on a diurnal timescale.

## 1 Introduction

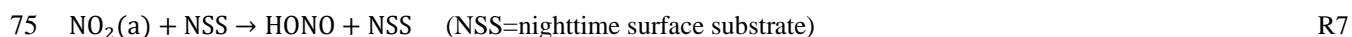
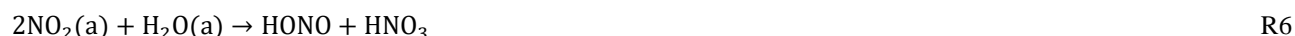
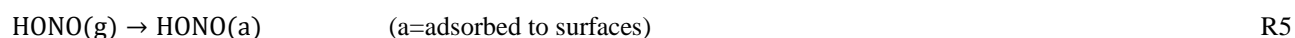
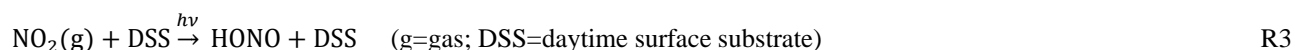
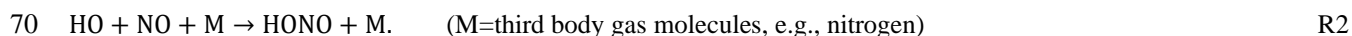
Vastly increased wildfire activity and intensity is a challenging issue in many parts of the world including the western United

30 States, and it is strongly linked to warming surface temperatures and earlier spring snowmelt (Westerling, 2016). Wildfire is a significant source of nitrogen oxides ( $\text{NO}_x = \text{NO} + \text{NO}_2$ ) and nitrous acid (HONO), as well as other important trace gases and particulate matter.  $\text{NO}_x$  serves as a key precursor to atmospheric ozone ( $\text{O}_3$ ) and secondary aerosols in the presence of organic compounds; in wildfire plumes  $\text{NO}_x$  can be a limiting factor to  $\text{O}_3$  production owing to high emission molar ratios of non-methane organic carbon (NMOC) to  $\text{NO}_x$  (Akagi et al., 2011; Jaffe and Briggs, 2012). HONO is a major daytime photolytic  
35 precursor of hydroxyl radical (OH) via R1 that determines the atmospheric oxidative capacity and therefore the lifetimes of many other species in the atmosphere. Wildfire emitted HONO supplies the majority of OH in the first few hours after smoke emission in the daytime, and it greatly counteracts reduced OH production from  $\text{O}_3$  photolysis caused by high particle loading reducing actinic flux (Jaffe and Briggs, 2012; Peng et al., 2020; Theys et al., 2020). Wildfire emitted  $\text{NO}_x$  and HONO not only greatly impact the atmospheric chemistry in local regions close to the fire, but also contribute significantly to the reactive  
40 nitrogen (RN) burden thousands of kilometres downwind via transport and RN cycling, especially when mixed with fossil fuel combustion emissions (Jaffe et al., 2013; McClure and Jaffe, 2018; Westerling et al., 2006; Westerling, 2016).



Despite their important impacts on air quality, climate, and human and ecosystem health, the budgets of wildfire-derived  $\text{NO}_x$  and HONO are poorly constrained due to limited field measurements, high reactivity, and large spatiotemporal heterogeneity.  
45 Bottom-up approaches rely on limited emission factor measurements with uncertainty in HONO sources and chemistry; top-down approaches (i.e., satellite observations) have limited sensitivity in the lower troposphere and boundary layer, and again, are limited by large uncertainties in HONO sources and chemistry to interpret the satellite measurements. Although gas phase reaction between OH and NO (R2) ubiquitously produces HONO, it is far from sufficient to explain the observed HONO levels in numerous studies given the fast photolysis during the day (Su et al., 2011). HONO, along with  $\text{NO}_x$ , can be directly emitted  
50 from various sources including vehicle exhaust, biomass burning (BB) and microbially-driven soil emissions. In addition, it has been proposed that HONO can be produced from other RN species (e.g.,  $\text{NO}_2$  and nitrate) via various heterogeneous pathways (Fig. 1). Major secondary HONO production pathways during the day include heterogeneous  $\text{NO}_2$  conversion on photoactive surfaces (R3) (Ammann et al., 1998; George et al., 2005; Stemmler et al., 2006), and heterogeneous photolysis of nitrate including particulate nitrate ( $\text{p-NO}_3^-$ ) and nitric acid ( $\text{HNO}_3$ ) via R4 (Ye et al., 2016; Zhou et al., 2011). In past studies,  
55 heterogeneous conversion of  $\text{NO}_2$  to HONO on photoactive surfaces such as organic surfaces (R3) has been proposed to explain a missing HONO source (Ammann et al., 1998; George et al., 2005; Stemmler et al., 2006; Wong et al., 2012). Organic surfaces exist in both aerosol particles and soils at the surface (e.g., humic acids) but there is major uncertainty associated with quantifying available surface area and the  $\text{NO}_2$  uptake coefficient. During the night, surface (soils and aerosols) uptake is the predominant sink for HONO (R5), and heterogeneous conversion of  $\text{NO}_2$  to HONO has been widely accepted as the major  
60 secondary HONO production source during the night (R6 and/or R7). Although it is clear that heterogeneous  $\text{NO}_2$  hydrolysis (R6) can be a major pathway for nighttime HONO production (Finlayson-Pitts et al., 2003), recent work has also shown

compelling evidence for faster HONO formation via reduction of NO<sub>2</sub> on inorganic surfaces (e.g., iron-bearing minerals) and organic surfaces (e.g., quinone-rich humic acid) in soils and particulate matter (R7) (Scharko et al., 2017; Kebede et al., 2016; Martins-Costa et al., 2020). While the emission sources and heterogeneous pathways were hypothetically used to account for missing HONO sources (Stemmler et al., 2006; Su et al., 2011; Ye et al., 2016; VandenBoer et al., 2014; Donaldson et al., 2014a; Kebede et al., 2016; Scharko et al., 2017), their relative importance is poorly quantified due to large uncertainties associated with emission heterogeneity, surface area and composition, environmental condition (day versus night, temperature, relative humidity), quantification of heterogeneous reaction rate and knowledge gaps in detailed mechanisms. As a result, the HONO budget in the atmospheric boundary layer remains poorly constrained.



Stable isotopes hold unique promise to provide rigorous constraints on sources, chemical processing pathways and sinks of RN species, as they reflect isotopic signatures associated with these processes.  $\delta^{15}\text{N}$  ( $=[(^{15}\text{N}/^{14}\text{N})_{\text{sample}}/(^{15}\text{N}/^{14}\text{N})_{\text{air-N}_2}-1]\times 1000\%$ ) has shown great potential to trace atmospheric origin of NO<sub>x</sub> and its fate as nitrate (Hastings et al., 2009), whereas  $\delta^{18}\text{O}$  ( $=[(^{18}\text{O}/^{16}\text{O})_{\text{sample}}/(^{18}\text{O}/^{16}\text{O})_{\text{VSMOW}}-1]\times 1000\%$ ; VSMOW is Vienna Standard Mean Ocean Water) serves as a sensitive indicator for relative importance of major oxidants (i.e. O<sub>3</sub>, RO<sub>2</sub> and OH) that lead to NO<sub>x</sub> conversion (Thiemens, 2006). In particular, O<sub>3</sub> has an exclusively high  $\delta^{18}\text{O}$  as a result of mass-independent fractionation associated with its formation in the atmosphere and this anomaly is transferred to oxidized products such as NO<sub>2</sub>, HONO and HNO<sub>3</sub> (Thiemens, 2006).

Using our recently developed and validated sampling techniques in combination with offline isotopic composition analyses (Chai et al., 2019), we characterized for the first time  $\delta^{15}\text{N}$  of NO<sub>x</sub> and HONO as well as  $\delta^{18}\text{O}$ -HONO in ground-level wildfire plumes in the western US as part of two major field campaigns: Western Wildfire Experiment for Cloud Chemistry, Aerosol Absorption and Nitrogen (WE-CAN) in summer 2018, and Fire Influence on Regional to Global Environments Experiment – Air Quality (FIREX-AQ) in summer 2019. Here we report our findings based on samples collected in a mobile laboratory platform from three different wildfires; Rabbit Foot Fire (RF) in eastern Idaho, Williams Flats Fire (WF) in central Washington and Nethker Fire (NF) in northern Idaho (Fig. S1 and Fig. S2 in the supplement). Surface-based mobile sampling allowed us

90 to characterize young nighttime (YN), young daytime (YD), mixed daytime (MD), aged nighttime (AN), and aged daytime  
(AD) smoke. Physical smoke age determination using meteorological parameters near the ground is challenging due to large  
variations of wind speed and direction. Proxies involving total RN, NO<sub>y</sub> and ammonia (NH<sub>3</sub>) relative to carbon monoxide (CO)  
can only be used to qualitatively evaluate smoke age due to large uncertainties of source emission factors and complexity  
caused by photochemistry (Selimovic et al., 2019; Kleinman et al., 2007). In contrast, the concentration ratio between PM<sub>2.5</sub>  
95 and CO (PM<sub>2.5</sub>/CO) has shown potential for estimating smoke age (Yokelson et al., 2009; Selimovic et al., 2020). In this work,  
we determined the smoke conditions (young vs aged) primarily by comparing the field  $\delta^{18}\text{O}$ -HONO results with those obtained  
in our previous lab study that represents fresh emissions, with additional evaluation involving  $\delta^{15}\text{N}$ -HONO and relative  
concentration of HONO and NO<sub>2</sub> (Fig. 2). Note young and aged smoke refers to negligible and large proportion of secondarily  
produced HONO respectively. We also take into account smoke sampling locations (i.e., approximate distance from the  
100 wildfire) to confirm the smoke age estimate. In brief, largely elevated  $\delta^{18}\text{O}$ -HONO in field samples compared with those from  
the lab-controlled fires signifies significant atmospheric processing, and this will be discussed in detail below. Our grouping  
method using  $\delta^{18}\text{O}$ -HONO shows fairly consistent results with that suggested by PM<sub>2.5</sub>/CO for WFF and NF fire plumes  
(Kaspari et al., 2021). In addition to distinguishing aged smoke from young smoke, the grouped  $\delta^{18}\text{O}$  and  $\delta^{15}\text{N}$  also allow us  
to characterize potential mechanisms of secondary HONO formation in the aged smoke as well as NO-to-NO<sub>2</sub> oxidation  
105 pathways, with the HONO budget evaluated using the synergistic measurement of HONO, NO and NO<sub>2</sub> concentrations in the  
field.

## 2 Methodology

### 2.1 Description of Mobile Laboratory platforms: Molab and MACH-2

During the WE-CAN campaign in August of 2018, we conducted our measurements and sampling using the NOAA Chemical  
110 Science Division mobile laboratory (Molab), which was a cargo van with all instruments mounted on it. Meteorological  
instrumentation on the roof of the Molab provides temperature, relative humidity, wind speed, wind direction, altitude, and  
GPS coordinates. All additional instruments were mounted to the interior floor and ambient air is sampled through 1- or 2-  
meter Teflon inlets that exit the roof of the Molab via bored holes. NO and NO<sub>x</sub> concentrations were measured with a Thermo  
Scientific Model 42i chemiluminescence NO/NO<sub>x</sub> analyzer owned by Brown University, with  $\pm 0.4$  ppbv precision and 0.2  
115 ppbv zero noise at 1 minute time resolution. Note the NO<sub>x</sub> concentration measured using the chemiluminescence analyzer can  
be falsely elevated due to known interferences from NO<sub>y</sub> species, e.g., HONO and PAN. However, these data provide an upper  
limit of NO<sub>x</sub> level that supports the isotopic collections of NO<sub>x</sub>, HONO and nitrate. HONO and HNO<sub>3</sub> concentrations were  
measured using University of New Hampshire's dual mist chamber/ion chromatograph (MC/IC) system with an uncertainty  
of 3% at 5-minute resolution (Chai et al., 2019; Scheuer et al., 2003). During the FIREX-AQ field campaign in July-August

120 of 2019, we mounted our sampling instruments in the NASA Langley mobile aerosol characterization platform (MACH-2)  
(Kaspari et al., 2021).

## 2.2 Description sampling location and strategy

125 While our sampling strategy was similar in both years, the actual sampling approach differed in response to fire condition and  
accessibility to fresh smoke from the mobile platforms.

During the 2018 WE-CAN campaign, our ground measurement and sampling targeted smoke from Rabbit foot fire (RF) in the  
Challis area of Salmon-Challis National Forest in central Idaho, from August 9<sup>th</sup> to August 18<sup>th</sup> 2018 (Salmon-Challis National  
130 Forest, 2018). Measurements were made at various locations around the Challis area of Idaho impacted by the RF fire,  
consisting of 4 different conditions: young smoke during nighttime (YN), young smoke during daytime (YD), aged smoke  
during nighttime (AN), aged smoke during daytime (AD), and mixed daytime smoke (M) that contains smoke contributed by  
either night smoke or fresh smoke. To sample the young smoke, we drove the Molab to Morgan Creek Road (MCR), which  
extends into a valley that was several kilometres away from the edge of the fire. We observed heavy smoke that based on  
135 distance and wind speed was expected to transport from the RF fire burning locations to the valley within a few hours or less.  
Three night-trips and two day-trips were made to MCR. While the nighttime measurements were conducted while driving the  
daytime work was carried out while parked at a spike camp (i.e. a campsite for firefighters and support personnel) at the upper  
end of MCR; the spike camp was about two kilometres from the fire, which we were able to see while conducting the  
measurements. The aged smoke was sampled at three stationary sites located around the Challis area, each less than 30 km  
140 away from the RF fire. All of these sites were recreational vehicle parks that allowed for power plugins. A total of 7 nights  
and 4 days were measured. The sampling locations and driving map are shown in Fig. S1 with detailed information on the  
measurements listed in Table 1.

During 2019 FIREX-AQ, we investigated five wildfires in the western US including Shady fire (Idaho), Black Diamond fire  
145 (Montana), Williams Flats fire (Washington), Nethker fire (Idaho), and Little Bear fire (Utah) from Jul 24 to Aug 22 of 2019.  
We intensively sampled the emissions from Williams Flats fire and Nethker fire based on the large size and easy access to  
sampling locations (Fig. S2). Similar to the 2018 field campaign, the measurements were conducted under YN, YD, AN, AD  
and M conditions.

## 150 2.3 Collection of HONO, NO<sub>x</sub> and nitrate for isotopic analysis

Nitrogen oxides (NO<sub>x</sub> = NO + NO<sub>2</sub>), nitrous acid (HONO), particulate nitrate (p-NO<sub>3</sub><sup>-</sup>) and nitric acid (HNO<sub>3</sub>) were captured  
in the field using recently developed methods and sent to Brown University for analyses of isotopic composition (Chai et al.,

2019; Fibiger and Hastings, 2016; Chai and Hastings, 2018; Fibiger et al., 2014). In brief, HONO was completely captured at  
155 a pumping flow rate of  $\sim 10$  L/min with an annular denuder system (ADS), comprised (in order) of a Teflon particulate filter  
to remove  $p\text{-NO}_3^-$ , a Nylasorb filter to remove  $\text{HNO}_3$ , followed by two annular denuders, each coated with a premixed  
 $\text{Na}_2\text{CO}_3$ -glycerol-methanol- $\text{H}_2\text{O}$  solution following a standard EPA method (Chai and Hastings, 2018). Within 24 hours after  
each collection, the coating was extracted in 10 mL of ultrapure water (18.2 M $\Omega$ ) in two sequential 5 mL extractions. Particulate  
nitrate on the upstream Millipore filters and  $\text{HNO}_3$  from the Nylasorb filters, if there was any, were extracted by sonicating  
160 the filters in  $\sim 30$  mL ultrapure  $\text{H}_2\text{O}$  (18.2 M $\Omega$ ). Samples with  $[\text{NO}_3^-] > 1$   $\mu\text{M}$  were analyzed for isotopic composition  
(concentration techniques detailed below).

The denuder extracted solution with a pH of  $\sim 10$  was frozen and transported to Brown University for concentration and isotopic  
analysis, which was completed within 2 months after the sampling. The timescales for sample extraction and isotopic analysis  
165 preserve both the solution concentration and isotopic composition of HONO in the form of nitrite (Chai and Hastings, 2018).  
The two-denuder set up allows for minimization of the interference for both concentration and isotopic analysis from other N-  
containing species that could be trapped and form nitrite in residual amounts on the denuders, especially  $\text{NO}_2$ . Note HONO  
levels were above the minimum detection limit (0.07  $\mu\text{M}$  in extraction solution) and the breakthrough amount of HONO  
threshold is far from being reached given the concentrations (Table 1), flow rate ( $\sim 8$  L/min) and collection times. Isotopic  
170 analysis of nitrite required collection of a minimum amount of 10 nmol.  $\text{NO}_x$  was completely collected in an impinging solution  
containing 0.25 M  $\text{KMnO}_4$  and 0.5 M  $\text{NaOH}$  which oxidizes  $\text{NO}$  and  $\text{NO}_2$  to  $\text{NO}_3^-$  by pumping sampled air through a gas  
washing bottle at a flow rate of  $\sim 4$  L/min. Collection time for HONO ranged from 2-12 hours and that for  $\text{NO}_x$  ranged from  
0.75 – 2.5 hours depending on their mixing ratios to make sure sufficient samples were captured against blanks for isotopic  
analysis (Fibiger et al., 2014; Fibiger and Hastings, 2016; Wojtal et al., 2016). Particulate filters and Nylasorb filters were  
175 collected over 7-12 hours due to low concentration of particulate nitrate and  $\text{HNO}_3$ .

The samples from each collection system were retrieved and processed following the procedures described in Chai et al. (2019).  
All treated samples from  $\text{NO}_x$ , HONO,  $p\text{-NO}_3^-$  and  $\text{HNO}_3$  collection and their corresponding blanks were analyzed offline for  
concentrations of  $\text{NO}_2^-$  and  $\text{NO}_3^-$  with a WestCo SmartChem 200 Discrete Analyzer colorimetric system. The reproducibility  
180 of the concentration measurements was  $\pm 0.3$   $\mu\text{mol L}^{-1}$  ( $1\sigma$ ) for  $\text{NO}_2^-$  and  $\pm 0.4$   $\mu\text{mol L}^{-1}$  for  $\text{NO}_3^-$  when a sample was repeatedly  
measured ( $n = 30$ ). A detection limit of 0.07  $\mu\text{mol L}^{-1}$  for  $\text{NO}_2^-$  and 0.1  $\mu\text{mol L}^{-1}$  for  $\text{NO}_3^-$  was determined, and no detectable  
nitrite or nitrate was found in the blank denuder coating solution, whereas blank  $\text{NO}_3^-$  concentrations of  $\sim 5$   $\mu\text{M}$  are typical for  
the  $\text{NO}_x$  collection method (Fibiger et al., 2014; Wojtal et al., 2016). We only report the samples whose concentrations were  
at least 30% above  $\text{NO}_3^-$  present in the blank  $\text{KMnO}_4$  solution upon purchase to avoid increasing the error associated with the  
185 isotopic composition (Fibiger et al., 2014). Note that  $\text{NO}_3^-$  concentration was measured on the ADS solutions to verify whether  
and to what extent  $\text{NO}_2^-$  was oxidized to  $\text{NO}_3^-$  on denuder walls because the denitrifier method will convert both  $\text{NO}_3^-$  and  
 $\text{NO}_2^-$  to  $\text{N}_2\text{O}$  for isotopic analysis (see below).

190 Note that complete collection of HONO and NO<sub>x</sub> have been verified in various environments including biomass burning  
emissions. During the FIREX fire lab experiment, we applied the same method to quantify the HONO and NO<sub>x</sub> isotopic  
composition (Chai et al., 2019). The concentrations of HONO captured with our annular denuder system (ADS) well compared  
with 4 other high time resolution concentration measurement techniques, including mist chamber/ion chromatography  
(MC/IC), open-path Fourier transform infrared spectroscopy (OP-FTIR), cavity enhanced spectroscopy (CES), and proton-  
195 transfer-reaction time-of-flight mass spectrometer (PTR-ToF). In the same work, the NO<sub>x</sub> concentrations collected in the  
permanganate impinger was verified by real-time measurement with a chemiluminescence NO<sub>x</sub> analyzer. In addition, our NO<sub>x</sub>  
collection technique has been verified with real-time NO<sub>x</sub> concentrations in on-road, near-road and urban background  
environments (Wojtal et al., 2016; Miller et al., 2017). These agreements verify complete capture of HONO and NO<sub>x</sub> associated  
with biomass burning emissions using our techniques, which preserve the isotopic signatures without isotopic fractionation  
during the sampling process.

200

## 2.4 Isotopic analysis

The denitrifier method was used to complete nitrogen (<sup>15</sup>N/<sup>14</sup>N) and oxygen (<sup>18</sup>O/<sup>16</sup>O) isotope analyses of separate NO<sub>3</sub><sup>-</sup>  
samples converted from NO<sub>x</sub>, and NO<sub>2</sub><sup>-</sup> samples converted from HONO by quantitative conversion to N<sub>2</sub>O by denitrifying  
205 bacteria *P. aureofaciens* (Casciotti et al., 2002; Sigman et al., 2001). The isotopic composition of N<sub>2</sub>O is then determined by  
a Thermo Finnegan Delta V Plus isotope ratio mass spectrometer at *m/z* 44, 45 and 46 for <sup>14</sup>N<sup>14</sup>N<sup>16</sup>O, <sup>14</sup>N<sup>15</sup>N<sup>16</sup>O and <sup>14</sup>N<sup>14</sup>N<sup>18</sup>O,  
respectively. Sample analyses were corrected against replicate measurements of the NO<sub>3</sub><sup>-</sup> isotopic reference materials USGS34,  
USGS35, and IAEA-NO-3 (Böhlke et al., 2003), and that of the NO<sub>2</sub><sup>-</sup> isotopic reference materials N7373 and N10219.  
Precisions for δ<sup>15</sup>N-NO<sub>x</sub>, δ<sup>15</sup>N-HONO and δ<sup>18</sup>O-HONO isotopic analysis across each of the entire methods are ±1.3‰, ±0.6‰  
210 and ±0.5‰ respectively (Chai and Hastings, 2018; Fibiger et al., 2014).

## 3 Results and Discussion

### 3.1 Concentrations of HONO and NO<sub>x</sub>

Among the three fires, increased HONO concentrations were observed in young smoke during both night (0.2 – 2.0 ppbv) and  
day (2.5 ppbv), while HONO level is significantly lower in aged smoke during both night (0.06 – 1.0 ppbv) and day (0.05 –  
215 0.6 ppbv) as shown in Fig. 2 (a). Although median values show young night and day are significantly higher than aged smoke  
day and night, there is significant overlap between young nighttime and aged day and night for WF and Nethker fires. These  
ppbv to sub-ppbv HONO concentrations can be a major OH source in areas that are impacted by wildfire. We also determined  
the molar ratio HONO/NO<sub>2</sub> from the concentrations for each sample (Fig. 3), and the values represent the upper bound of

[HONO]/[NO<sub>x</sub>] (Table 1(a)). Median ratios of [HONO]/[NO<sub>2</sub>] for the five smoke conditions are 0.35 (YN), 0.12 (YD), 0.07 (AN), 0.09 (AD) and 0.04 (MD). The median ratios of [HONO]/[NO<sub>2</sub>] for the young smoke fall in the range of fresh emissions measured in the lab (0.13-0.53), and the field (0.05-0.33) (Yokelson et al., 2009; Selimovic et al., 2020 and references therein). Our results for YN are also in agreement with airborne measurements (0.34±0.08) from the BB-Flux campaign that occurred in parallel with WE-CAN, but are lower than the WE-CAN airborne observation of 0.72±0.34 during the day (Theys et al., 2020; Peng et al., 2020). It is worth noting that the majority of the WE-CAN airborne data overlap with the BB-Flux results and our measurements, and Peng et al. (2020) associate very high ratios with different transport dynamics of fresh plumes. The concentration results for the ADS collected [HONO] agree well with that measured via MC/IC in real-time and averaged over the ADS sampling periods (Fig. S3). The good agreement between these techniques sampling the same plumes near the ground, and previous agreement with other HONO and NO<sub>x</sub> observation methods suggest the concentrations are accurate (see also Section 2.3). It is important to also consider possible interference of peroxyacetyl nitrate (PAN) with NO<sub>x</sub> collected in the alkaline permanganate solution for biomass burning conditions (Jaffe and Briggs, 2012). There is minimum PAN formed in fresh biomass burning (BB) emissions and young smoke of less than half an hour, based upon previous lab and field measurements, as well as modeling studies (Stockwell et al., 2014; Yokelson et al., 2009; Alvarado et al., 2010, 2015). In aged BB plumes in the upper troposphere, PAN can form rapidly at low temperatures and act as a temporary NO<sub>x</sub> reservoir, reaching a maximum PAN/NO<sub>y</sub> ratio of 0.3 (comparable to NO<sub>x</sub>/NO<sub>y</sub>) within ~2 to 4 hours of aging after emission (Yokelson et al., 2009; Liu et al., 2016; Akagi et al., 2012). Though we note that these results are all from airborne measurements. There are no ground-level measurements for PAN in BB plumes during WE-CAN or FIREX-AQ, nor from other field studies, to the best of our knowledge. PAN is thermally unstable in the boundary layer during summertime, and its main loss process in the atmosphere is thermal decomposition to release NO<sub>2</sub>. The lifetime of PAN is on the order of 1 hour or less at 20 °C and above (Talukdar et al., 1995; Fischer et al., 2010). We therefore expected PAN in near-ground air to maintain low levels or less due to photochemistry and thermal decomposition. Thus, given the short lifetime and the sample integration time of over 40 min to 2 hours' timescale, PAN is unlikely to interfere with our NO<sub>x</sub> results.

In the aged smoke, [HONO]/[NO<sub>2</sub>] are greatly reduced to median values of 0.05 and 0.07 observed for AN and AD, respectively, lower than the lab-derived range (Fig. 3). Wildfire plumes near ground-level are expected to be more diluted than that directly injected upwards during the day and the loss of HONO due to photolysis in the plume and/or surface reactions would be expected to be much faster than that in higher altitude dense plumes. The very low ratios indicate that HONO was lost faster than NO<sub>x</sub>; however, given the 10-20 minute lifetime of HONO against photolysis during the day and up to a couple of hours during the night (Nie et al., 2015), and considering aged smoke was sampled 10s of km from the fire, HONO levels may be maintained via secondary chemistry due to the high particle loadings and other terrestrial surface reactions in wildfire plumes (Alvarado and Prinn, 2009). While the concentration data are valuable for the ground-based setting near the fires, considerable uncertainty exists in the rate coefficients of the heterogeneous processes in daytime, as well as the HONO and NO<sub>2</sub> uptake coefficient and surface area densities (Appendix A). This makes it challenging to quantify the relative contribution



of each potential pathway to the observed HONO budget.

### 3.2 Isotopic signatures of HONO and NO<sub>x</sub>

In the 2016 FIREX Fire Laboratory experiment, we obtained  $\delta^{15}\text{N}$  of NO<sub>x</sub> and HONO, as well as  $\delta^{18}\text{O}$  of HONO in direct  
255 emissions from controlled burning of various vegetation biomass representative of the western US (Chai et al., 2019). The lab-  
based  $\delta^{15}\text{N}$  and  $\delta^{18}\text{O}$  results serve as source signatures of biomass burning (BB) emissions:  $\delta^{15}\text{N-NO}_x$  (-4.3‰ to +7.0‰) and  
 $\delta^{15}\text{N-HONO}$  (-5.3‰ to +5.8‰) are derived from biomass N and the transformation in the combustion process, and  $\delta^{18}\text{O-}$   
HONO (5.2‰ to 15.2‰) incorporates  $\delta^{18}\text{O}$  of molecular oxygen and water via combustion reactions (Chai et al., 2019). In the  
field, we expect that once NO<sub>x</sub> and HONO are released and transported, atmospheric processing including photochemistry and  
260 nighttime chemistry would cause the isotopic composition of emitted NO<sub>x</sub> and HONO to change.

By directly comparing the field-measured  $\delta^{18}\text{O-HONO}$  with that measured from lab-controlled burning, we separate the data  
observed in young smoke from that in aged smoke. Very young smoke largely reflects fresh wildfire emissions without  
significant atmospheric processing, while aged smoke  $\delta^{18}\text{O-HONO}$  should deviate from the lab values due to the influence of  
secondary chemistry involving RN cycling. The  $\delta^{18}\text{O-HONO}$  of young nighttime smoke ranged from 4.8‰ to 32.3‰ with a  
265 median value of 19.0‰, while the value in a single young daytime sample was 25.6‰ (Fig. 2(b)). There is a major overlap  
between the lab results and young nighttime smoke, but with some higher  $\delta^{18}\text{O-HONO}$  values in the field observations. These  
results suggest the HONO sampled in young smoke was dominated by primary BB emissions from the nearby wildfire, but  
included contributions of secondarily produced HONO. By contrast,  $\delta^{18}\text{O-HONO}$  is greatly elevated in aged smoke from all  
three fires both day and night. In addition, two aged smoke samples are labelled as mixed smoke because the collection interval  
270 included both sunlit and dark periods. The enrichment of  $\delta^{18}\text{O-HONO}$  (up to 78‰), regardless of location and time, suggests  
that HONO in these conditions is produced by secondary chemistry involving NO, NO<sub>2</sub> and nitrate, which transfer high  $\delta^{18}\text{O}$   
values due to O<sub>3</sub> influence via photochemistry (Appendix B) (Thiemens, 2006; Michalski et al., 2003). The varying  $\delta^{18}\text{O-}$   
HONO values reflect different oxidizing environments, i.e., NO-to-NO<sub>2</sub> conversion via RO<sub>2</sub> versus O<sub>3</sub>. These branching ratios  
can be determined if we resolve the dominant pathways for HONO production.

275  $\delta^{15}\text{N-HONO}$  in the young smoke ranges from -0.3‰ to +7.4‰ with a median value of 2.8‰ for YN, and +3.4‰ for YD,  
whereas that in the aged smoke shows decreased median values -2.9‰ and -1.8‰ for AN and AD respectively. In addition,  
the daytime aged smoke exhibits the largest variability (Fig. 2(c)), and this likely reflects daytime HONO secondary chemistry.  
It is noted that  $\delta^{15}\text{N-NO}_x$  and  $\delta^{15}\text{N-HONO}$  measured across the entire period of all three fires at ground level ranges from -  
4.3‰ to +8.7‰ and -6.7‰ to +7.4‰, respectively, with the majority overlapping with the corresponding ranges found in the  
280 Fire Laboratory experiment and no significant difference in mean values ( $p$ -value >0.5) (Chai et al., 2019; Fibiger and Hastings,  
2016). This consistency suggests  $\delta^{15}\text{N}$  is a reliable tracker generally for BB derived NO<sub>x</sub> and HONO, although there is clear  
variability between the different smoke conditions that can refine our understanding of reactive N cycling. We note again that,

285 although no near-ground PAN measurements in BB plumes are available, the isotopic results also suggest that PAN interference is not important to the  $\delta^{15}\text{N-NO}_x$  results. For aged smoke, we would expect  $\delta^{15}\text{N-NO}_x$  to decrease from that in fresh emissions due to partial transformation of  $\text{NO}_x$  to additional oxidized N products (e.g., PAN), as well as isotopic exchange between  $\text{NO}_x$  and these oxidized species; both processes will leave  $^{15}\text{N}$  depleted in  $\text{NO}_x$  and  $^{15}\text{N}$  enriched in PAN (Walters and Michalski, 2015). If PAN existed at significant concentrations that were 1) comparable with  $\text{NO}_x$  in the atmosphere, and 2) completely collected in the permanganate solution, then the  $\delta^{15}\text{N}$  would reflect the overall  $\delta^{15}\text{N}$  of  $\text{NO}_x + \text{PAN}$  in the final reduced permanganate solution. In this case, we would expect that aged smoke would not shift from the  $\delta^{15}\text{N-NO}_x$  range of young smoke, because  $\delta^{15}\text{N}$  shifts in both PAN and  $\text{NO}_x$  could offset each other. However, our observed  $\delta^{15}\text{N-NO}_x$  mean values for both aged daytime and nighttime smoke are significantly ( $p < 0.05$ ) lower than that of the young smoke, a good indicator of a lack of PAN interference on the isotopic results (see also Miller et al. (2017)).

Our prior lab-controlled burning study revealed a linear relationship between  $\delta^{15}\text{N-HONO}$  and  $\delta^{15}\text{N-NO}_x$ , with  $\delta^{15}\text{N-HONO}$  slightly more negative than  $\delta^{15}\text{N-NO}_x$  in fresh BB emissions (Chai et al., 2019). This  $\delta^{15}\text{N}$  relationship is plotted as a solid line together with all field observations to illustrate the potential influence of atmospheric processing on the  $\delta^{15}\text{N-HONO}$  and  $\text{-NO}_x$  (Fig. 4). The plot can be sub-divided into three regimes. In regime I, we find all of the  $\delta^{15}\text{N}$  of  $\text{NO}_x$  and HONO in young smoke from both daytime and nighttime. In this young smoke regime, more positive  $\delta^{15}\text{N}$  than that of the rest of our samples is found for both species and all samples concur with the  $\delta^{15}\text{N}$  relationship found for fresh emissions (Fig. 4). This, along with the low  $\delta^{18}\text{O-HONO}$  associated with these samples (Fig. 2(b)), confirms HONO is not significantly affected by secondary chemical processing in the air mass captured from fresh smoke. Regime II is filled with the results of daytime aged smoke ~30 km away from the RF fire; these results exhibited much more positive  $\delta^{15}\text{N-HONO}$  than  $\delta^{15}\text{N-NO}_x$  by 3‰-6‰, and the largest (positive) discrepancy from the BB  $\delta^{15}\text{N}$  relationship line, as shown in the upper left region of Fig. 4. The daytime aged smoke also exhibited the highest values of  $\delta^{18}\text{O-HONO}$  observed (Fig 2.). All samples of aged nighttime smoke that were collected fall in regime III. While the majority of the regime III data fall within the 95% confidence interval for the lab-based  $\delta^{15}\text{N}$  relationship, there is a tendency for these samples to have  $\delta^{15}\text{N-HONO}$  that was more negative than  $\delta^{15}\text{N-NO}_x$  to different degrees of up to -8.7‰. In particular, we hypothesize that the combination of more negative  $\delta^{15}\text{N-HONO}$  values and elevated  $\delta^{18}\text{O-HONO}$  indicate secondary production of HONO. We next explore quantitative use of  $\delta^{15}\text{N-NO}_x$ ,  $\delta^{15}\text{N-HONO}$  and  $\delta^{18}\text{O-HONO}$  to understand the isotopic shifts in terms of secondary chemistry involving RN cycling.

### 3.3 Isotopic mass balance modelling

310 In aged smoke, the observed  $\delta^{18}\text{O-HONO}$  enhancement and shift of  $\delta^{15}\text{N}$  values away from the  $\delta^{15}\text{N NO}_x\text{-HONO}$  line, as a result of RN cycling, would be expected to derive from the integrated kinetic isotopic fractionation (expressed as enrichment factor  $^{18}\epsilon$  and  $^{15}\epsilon$ ) associated with each of the loss/production processes (Fig. 1) weighted by their relative contribution to the budget. By definition,  $\epsilon = (\alpha - 1) \times 1000\text{‰}$ , with fractionation factor  $\alpha$  referring to the rate coefficient ratio between the heavy

isotopologue and the light isotopologue. For  $\delta^{18}\text{O}$ -HONO, we also took into account transferring effect of oxygen from  
 315 different O-containing reactants that produce HONO (as explained in Appendix B). In order to elucidate the relative role each  
 process plays in the HONO budget, we constructed an isotopic mass balance model for  $\delta^{15}\text{N}$  and for  $\delta^{18}\text{O}$ .

In aged smoke, a deviation in  $\delta^{15}\text{N}$ , represented as  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  ( $=\delta^{15}\text{N-HONO} - \delta^{15}\text{N-NO}_x$ ), is simulated following Eq. (1),  
 where  $f$  is the fraction of reaction  $i$  (R numbering in Section 1) to total loss (L) or production (P) of HONO.  $\delta^{18}\text{O}$ -HONO is  
 320 simulated following Eq. (2), where the change of  $\delta^{18}\text{O-HONO}_{i,P}$  arises from, in addition to kinetic isotopic fractionation, the  
 transferring of  $\delta^{18}\text{O}_{i,t}$  (Eq. (3)) in the reactant (OH, NO,  $\text{NO}_2$ ,  $\text{H}_2\text{O}$ , and  $\text{NO}_3^-$ ) to the product HONO, as HONO contains two  
 O atoms that may stem from more than one reactant (Appendix B).  $\delta^{18}\text{O}$  of all possible reactions that produce HONO are  
 evaluated as tabulated in Table S1 in the supplement, to help determine  $\delta^{18}\text{O}$  of NO,  $\text{NO}_2$  and HONO. The isotopic enrichment  
 factors  $^{15}\epsilon$  and  $^{18}\epsilon$  associated with each of the reactions R1-R7 illustrated in Fig. 1 are computed via theoretical principles, as  
 325 none of these key parameters are currently available in literature (Appendix B).

$$\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x} = \sum_{i,L} (f_{i,L} \times \Delta\delta^{15}\text{N}_{i,L}) + \sum_{i,P} (f_{i,P} \times \Delta\delta^{15}\text{N}_{i,P}) \quad (1)$$

$$\delta^{18}\text{O-HONO} = \sum_{i,L} (f_{i,L} \times ^{18}\epsilon_{i,L}) + \sum_{i,P} (f_{i,P} \times \Delta\delta^{18}\text{O-HONO}_{i,P}) \quad (2)$$

$$\Delta\delta^{18}\text{O-HONO}_{i,P} = \delta^{18}\text{O}_{i,t} + ^{18}\epsilon_{i,P} \quad (3)$$

### 330 3.3.1 Modelling of $\delta^{15}\text{N}$ of HONO and $\text{NO}_x$ in aged daytime and nighttime smoke

We first simulated  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  for both daytime and nighttime aged condition using this model. According to the potential  
 HONO- $\text{NO}_x$  chemistry in ground areas impacted by wildfire smoke plumes (Fig. 1), HONO is expected to be predominantly  
 lost to photolysis (R1) during the day. It is well known that HONO can be produced via gas-phase radical recombination  
 reaction between NO and OH (R2) (Platt et al., 1980). However, the rate of R2, calculated from the rate coefficient, the typical  
 335 daytime OH concentration ( $1-2 \times 10^6$  molecule  $\text{cm}^{-3}$ ) (de Gouw et al., 2006) in biomass burning plumes and our measured mean  
 $\text{NO}_x$  concentration, can only account for up to 15% of the HONO production rate (Appendix A and Table A1). Under a typical  
 pseudo steady state approximation ( $d[\text{HONO}]/dt \approx 0$ ), additional sources of HONO must be included to balance the HONO  
 budget. Thus, we modelled three scenarios varying the relative contribution of R2 as 5%, 10% and 15%. With rapid photolytic  
 loss, HONO has a lifetime nearly two orders of magnitude shorter than the lifetime of NO in R2 as well as that of  $\text{NO}_2$  in R3  
 340 and nitrate in R4 (Fig. 1); thus, the  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  is mostly sensitive to the change in  $\delta^{15}\text{N}_{\text{HONO}}$  immediately upon photolysis,  
 but overall remains constant associated with R2-R4 within the timescale of HONO photolysis. We quantify the remaining  
 HONO fraction from secondary production,  $f_{rp}$ , to represent HONO that has been produced but not yet photolyzed. Thus, the  
 daytime  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  for aged smoke was simulated as a function of  $f_{rp}$  following a Rayleigh-type isotopic fractionation  
 scheme (Fig. 5). Generally,  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  follows an exponential increase as  $f_{rp}$  decreases. In other words, as more photolysis

345 occurs the difference in the remaining  $\delta^{15}\text{N-HONO}$  and the  $\delta^{15}\text{N-NO}_x$  increases, and this is driven by the negative value of  $^{15}\epsilon_1$  which tends to enrich  $^{15}\text{N}$  in the HONO reactant (R1). The simulation was carried out for two different sets of HONO production mechanisms, with HONO photolysis being the dominant loss fate. With mechanism M1 (solid lines in Fig. 5), photo-induced surface  $\text{NO}_2$ -to-HONO conversion (R3) is the major pathway in addition to gas-phase  $\text{OH}+\text{NO}$  (R2) to produce HONO. As  $^{15}\epsilon_2$  has a positive value, larger R2 contribution leads to higher  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$ . With mechanism M2 (dashed line  
350 in Fig. 5), nitrate photolysis (R4) is included in addition to R2 and R3 in the HONO production mechanism. Taking the contribution of R2 of 10% as a constant, three scenarios were modelled by varying the relative contribution of R3 (75%-85%) and R4 (5%-15%). The results suggest larger R4 contribution yields lower  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  due to severe  $^{15}\text{N}$  depletion associated with nitrate photolysis ( $^{15}\epsilon_4 \leq -47.9\%$ ) (Appendix B). Importantly, the addition of R4 in M2 also lowers  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  compared to M1. By applying the field-observed  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  for the aged daytime smoke to the model, we  
355 solved  $f_{\text{TP}}$  for all scenarios and plotted these as circles in Fig. 5. All five daytime aged data from RF can be reproduced by M1 under all three scenarios; by contrast, via M2, none of the three scenarios can explain the two highest  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  observed in the field. As such, we conclude R4 plays a minor role (<5%) in the secondary HONO production in the aged daytime smoke during our sampling periods. Rather, HONO forms primarily via R2 and R3 during the day in the areas impacted by aged wildfire smoke. However, there are two limitations to the modeling results. First, as the  $^{15}\text{N}/^{14}\text{N}$  fractionation associated with  
360 R3, R6 and R7 are not distinguishable with our current parameterization (Appendix B.1.2 and B.2.2), we cannot rule out the potential importance of heterogeneous  $\text{NO}_2$ -to-HONO conversions (R6 and R7) in daytime. Second, it should be noted that the results represent our best estimate of the average relative importance of R2-R4 for HONO production during our HONO sampling periods (2-10 hours) for the aged daytime plume. Due to the long sample integration time, our samples were influenced by both aged smoke and near-background air when the smoke was very diluted. Under the  $\text{NO}_x$ -limited condition  
365 (low  $\text{NO}_x < 1$  ppbv) in remote background air, nitrate photolysis is expected to be the major secondary HONO source (Ye et al., 2016; Zhou et al., 2011), which cannot be ruled out by our results. Isotopic measurement techniques with higher time resolution will be required to achieve real-time quantification of the HONO budget.

For the nighttime smoke, we simulated that the HONO budget is maintained by R5-R7 (Fig. 1).  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  reflects the combination of kinetic isotopic fractionation  $^{15}\epsilon_5$  associated with the HONO loss R5 and production reactions (R6 and R7 in  
370 proportion). With our calculated uptake  $^{15}\epsilon_5$  (-2‰), and estimated  $^{15}\epsilon_6$  or  $^{15}\epsilon_7$  (ranging from -2.9‰ to -4.5‰), we obtained  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  ranging from -0.9 to -2.5‰ when uptake and production occurs at a similar time scale (rate coefficient), and this can explain the majority of observed aged nighttime results (regime III, Fig. 4). Two aged nighttime points sampled for RF (Aug 16 and 17, 2018) fall outside of the predicted range, with much lower  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  (-8.7‰ and -5.5‰ respectively). These two samples were associated with 2-10 times elevated  $\text{NO}_x$  concentration compared to the previous 4 nights and likely  
375 higher concentrations of particulate matter (Fig 2(a); Fig. S4 in the supplement). This could cause an accelerated conversion of  $\text{NO}_2$ -to-HONO, which is not accounted for in the steady state estimation above, leading to the much lower  $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$  values that were observed.

### 3.3.2 Modelling of $\delta^{18}\text{O}$ of HONO in aged daytime and nighttime smoke

380  $\delta^{18}\text{O}$ -HONO of daytime aged smoke was modelled following M1 (R1-R3) derived based upon the  $\delta^{15}\text{N}$  modelling results: NO and  $\text{NO}_2$  are cycled via  $\text{NO}_2$  photolysis and NO oxidation by  $\text{O}_3$  and/or peroxy radicals ( $\text{RO}_2$  including  $\text{HO}_2$ ) during the day, through which  $\delta^{18}\text{O}$  of  $\text{O}_3$  and  $\text{RO}_2$  can be passed to NO and  $\text{NO}_2$  via mass transfer (Eqs. (B9)-(B11)).  $\text{O}_3$  is known to have an intrinsically high  $\delta^{18}\text{O}$  value of up to  $\sim 117\%$  caused by unique isotopic fractionation associated with photochemical gas-phase  $\text{O}_3$  formation (Thiemens, 2006), while OH and  $\text{RO}_2$  have very low  $\delta^{18}\text{O}$  values (Thiemens, 2006).  $\text{O}_3$  participation in reactive N cycling involving  $\text{NO}_x$  (R8) results in high  $\delta^{18}\text{O}$  of  $\text{NO}_2$  (Michalski et al., 2003; Walters et al., 2018). In pseudo  
385 photochemical steady state, NO and  $\text{NO}_2$  are expected to have similar  $\delta^{18}\text{O}$  that is a result of competition between  $\text{O}_3$  and  $\text{RO}_2$  oxidation (R8-R10), expressed as  $f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}}$  via Eqs. (4) and (5) below.



$$\delta^{18}\text{O}\text{-NO} \approx \delta^{18}\text{O}\text{-NO}_2 = f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}} \times \delta^{18}\text{O}\text{-O}_3 + (1 - f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}}) \times \delta^{18}\text{O}\text{-RO}_2 \quad (4)$$

$$f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}} = \frac{k_{\text{NO}+\text{O}_3}[\text{O}_3]}{k_{\text{NO}+\text{O}_3}[\text{O}_3] + k_{\text{NO}+\text{RO}_2}[\text{RO}_2]} \quad (5)$$

395 The  $\delta^{18}\text{O}$  signature is subsequently passed to HONO when it is produced from NO (R2) and  $\text{NO}_2$  (R3) during the day and from  $\text{NO}_2$  (R6 and R7) during the night, and thus  $\delta^{18}\text{O}$ -HONO is a positive linear function of  $f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}}$  if kinetic isotopic fractionation ( $^{18}\epsilon$ ) associated with these processes are fixed values (as calculated in Appendix B). Given that HONO is predominantly produced via R2 and R3 in aged daytime smoke (Fig. 5),  $\delta^{18}\text{O}$ -HONO was simulated following the three M1 scenarios with the contribution of R2 varying from 5% to 15%. All three scenarios reproduced the range of our field results  
400 for aged daytime smoke, further pointing to M1 as explaining the HONO in this environment. In addition, the variation of  $\delta^{18}\text{O}$  was driven by differing oxidation that is determined by  $f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}}$ , which depends on the relative concentration of  $\text{O}_3$  to  $\text{RO}_2$  (Figs. S5 and S6).  $f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}}$  corresponding to each observed  $\delta^{18}\text{O}$ -HONO were solved and plotted in Fig. 6(a). We found  $f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}}$  decreased by less than 0.02 as the contribution of R2 to total HONO production decreased from 15 to 5%. On the other hand,  $\delta^{18}\text{O}$ -HONO changes sensitively with varying  $f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}}$ , increasing from 50.2‰ to 78.0‰ as the fraction of NO  
405 oxidized by  $\text{O}_3$  rather than  $\text{RO}_2$  increases from 0.34 to 0.65.

$\delta^{18}\text{O}$ -HONO of nighttime aged smoke was modelled following the nighttime chemistry (R5-R7), i.e. taking  $\text{NO}_2$  conversion as the source and surface uptake as the sink. In areas impacted by nighttime aged smoke, HONO forms from wildfire derived  $\text{NO}_2$  residing in the nocturnal boundary layer. As the two pathways (R6 and R7) for heterogeneous  $\text{NO}_2$  conversion lead to very different  $\delta^{18}\text{O}$ -HONO stemming from different  $\delta^{18}\text{O}$  transfer (Appendix B), we examined the relative importance of the two pathways for HONO production by varying the relative contribution between the two pathways and comparing to the observed  $\delta^{18}\text{O}$ -HONO (Fig. 6(b)). If HONO is constrained to exclusively form via R6 (surface hydrolysis), the model would require an unrealistic  $f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}} > 100\%$  to explain  $\delta^{18}\text{O}$ -HONO  $> 55\%$ . Even for samples with lower  $\delta^{18}\text{O}$ -HONO values (34%-52%), the high branching ratio  $f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}} (> 0.6)$  required to create such large enrichment is unrealistic for BB environments. In particular,  $[\text{O}_3]/[\text{RO}_2]$  converted from  $f_{\text{O}_3/\text{RO}_2}^{\text{NO}}$  solved under this mechanism is at least twice as large as values derived from the previous field measurement of aged wildfire smoke (Baylon et al., 2018). By contrast, inclusion of R7 in addition to R6 in rate ratios 3:1 and 20:1 based on previous lab studies (Kebede et al., 2016; Scharko et al., 2017) can elevate the modelled  $\delta^{18}\text{O}$ -HONO and explain all observed  $\delta^{18}\text{O}$ -HONO values. This suggests  $\text{NO}_2$ -to-HONO heterogeneous conversion catalysed by surface-hosted iron oxides and quinone (R7) in the nighttime aged smoke proceeds significantly faster than  $\text{NO}_2$  hydrolysis (R6). Our isotopic analyses provide evidence for participation of such pathway in BB environments, and also shows the capability to constrain the relative importance between these two pathways. Although the daytime  $\delta^{18}\text{O}$ -HONO can be larger than that of nighttime aged smoke, similar  $[\text{O}_3]/[\text{RO}_2]$  ratios are derived from our solved  $f_{\text{O}_3/(\text{O}_3+\text{RO}_2)}^{\text{NO}}$  and are consistent with the limited field measurements (Parrington et al., 2013; Baylon et al., 2018), and further indicate the important role peroxy radicals play as an oxidant in wildfire smoke impacted environments.

425

#### 4 Conclusion

As wildfire has enormously impacted climate, air quality and ecosystems in the past and is expected to worsen (Westerling, 2016), accurately tracking wildfire derived reactive nitrogen species (i.e.  $\text{NO}_x$  and HONO) and their cycling is extremely important for quantifying and mitigating key pollutants such as  $\text{O}_3$  in wildfire impacted areas both close to the fire and thousands of kilometres downwind. We show  $\delta^{15}\text{N}$ -HONO and  $\delta^{15}\text{N}$ - $\text{NO}_x$  can serve as a powerful tool to track BB sources and constrain secondary HONO production pathways. With the help of field-observed  $\delta^{18}\text{O}$ -HONO, we grouped our measured relationship between the  $\delta^{15}\text{N}$ -HONO and  $\delta^{15}\text{N}$ - $\text{NO}_x$  into three different regimes, which clearly distinguish among young wildfire plume, aged daytime plume as well as aged nighttime plume. The  $\delta^{15}\text{N}$  results allow for constraining the daytime HONO budget particularly secondary production mechanisms via the isotope mass balance simulation. The use of excess  $\delta^{15}\text{N}$  ( $\Delta\delta^{15}\text{N}_{\text{HONO-NO}_x}$ ) also provides an approach for constraining HONO budgets in other environmental settings, such as urban

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ambient areas and remote areas including forest and polar regions. Furthermore, by combining  $\delta^{15}\text{N}$  emission source signatures and chemical fractionation characteristics, we could potentially track the impact and relative role of wildfire derived reactive nitrogen more extensively when the plume transfers thousands of kilometres downwind and mixes with other air such as urban  
440 plumes. In addition, the  $\delta^{18}\text{O}$ -HONO results not only offer direct evidence for secondary production of HONO that allows for determination of the NO oxidizing branching ratio between  $\text{O}_3$  and  $\text{RO}_2$ , but also constrains nighttime HONO production mechanisms. We expect to apply the  $\delta^{18}\text{O}$ -HONO approach to a variety of atmospheric settings for constraining the HONO budget and its cycling with other reactive nitrogen species as well as  $\text{O}_3$ . As such, online isotopic measurement techniques with higher time resolution will benefit the use of stable isotopes and broaden its application in atmospheric chemistry. In the  
445 meantime, in order to more accurately quantify the relative contribution of these potential pathways, further experimental and theoretical investigations on isotopic fractionation characteristics of each pathway under various environmental conditions are required.

#### 450 **Appendix A. Overview of HONO budget quantification under different conditions based upon concentrations**

A common approach to quantitatively understand the wildfire-derived HONO budget—its direct emissions, secondary productions and sinks—is to use concentration-based mass balance calculation. Ideally, if we know the rate coefficients and reactant concentrations for each of the pathways, we would be able to quantify the relative contribution of each pathway to the total HONO concentration measured in the field under the assumption of pseudo steady-state approximation (PSSA) as  
455 described in Eq. (A1), where  $R_{\text{emission}}$ ,  $R_{\text{production}}$  and  $R_{\text{loss}}$  are rate of emission, production and loss respectively. In aged smoke, we expect HONO is almost exclusively produced from secondary formation. During the day, HONO is predominantly lost to photolysis with a coefficient depending on solar zenith angle differing with time of the day, while one or more reactions of R2-R4 may be responsible for producing HONO (Fig. 1). Under PSSA, using the well quantified rate coefficient  $k_2$ , observed NO and HONO concentrations, estimated OH concentration, and TUV model calculated HONO photolysis coefficient  $j_{\text{HONO}}$ ,  
460 we estimated the ratio of R2 to the total HONO production ( $P_{\text{OH+NO}}$ ) via Eq. (A2) and found R2 can only contribute 2%-15% (Table A1) of the total HONO production under the ambient conditions when the five aged-day samples were collected. This suggests at least 85% of HONO was produced from heterogeneous HONO formation via R3 and/or R4.

$$\frac{d[\text{HONO}]}{dt} = R_{\text{emission}} + R_{\text{production}} - R_{\text{loss}} \approx 0 \quad (\text{A1})$$

$$P_{\text{OH+NO}} = \frac{k_2[\text{OH}][\text{NO}]}{j_{\text{HONO}}[\text{HONO}]} \quad (\text{A2})$$

465 HONO production from photo-enhanced  $\text{NO}_2$  conversion has been proposed to take place on various types of surfaces. However, the uptake coefficient ( $\gamma_{\text{NO}_2 \rightarrow \text{HONO}}^{h\nu}$ ), which indicates the probability of  $\text{NO}_2$  collisions with a surface that results in formation of a HONO molecule, varies by at least three orders of magnitude depending on the specific type of surface materials.

For instance,  $\gamma_{NO_2 \rightarrow HONO}^{hv}$  on soot particles was found to range from  $3.7 \times 10^{-4}$  to  $1.1 \times 10^{-3} \text{ s}^{-1}$  (Ammann et al., 1998), while that on surfaces comprised of humic acid was measured as  $2-8 \times 10^{-5} \text{ s}^{-1}$  in several lab studies (Stemmler et al., 2006; Scharko et al., 2017). The latter is consistent with daytime modeling results of  $6 \times 10^{-5} \text{ s}^{-1}$  (Wong et al., 2013). Additionally, much smaller ( $10^{-7} - 10^{-6} \text{ s}^{-1}$ )  $\gamma_{NO_2 \rightarrow HONO}^{hv}$  was obtained for metal oxide surface such as  $TiO_2/SiO_2$  (Ndour et al., 2008).

Daytime photolysis of nitrate ( $HNO_3$  and  $pNO_3^-$ ) via R4 has also been proposed as an important renoxification pathway that produces HONO and  $NO_2$  in low  $NO_x$ /remote environments (Zhou et al., 2011) as well as high  $NO_x$ /urban settings with abundant urban grime (Baergen and Donaldson, 2016, 2013). The  $p-NO_3^-$  and surface-adsorbed  $HNO_3$  were found to be photolyzed with rate coefficients 2-3 orders of magnitude larger than gas-phase  $HNO_3$ , and possess lifetimes as low as a few hours (Ye et al., 2017). However, the rate coefficient of R4 is poorly constrained. Not only have the branching ratio between  $NO_x$  producing channel and HONO forming channel been poorly known (Baergen and Donaldson, 2016), but previous laboratory measured nitrate photolysis rate coefficients also vary by up to 3 orders of magnitude (Ye et al., 2017). The uncertainty is even greater because it is complicated by dependence on relative humidity, particle composition and pH.

During the night, HONO is primarily lost to uptake on surfaces including aerosols and soils and the uptake coefficient can be expressed by Eq. (A3)

$$L_{HONO}^{uptake} = 0.25 \times \gamma_{HONO} \times \overline{\omega_{HONO}} \times S/V \times 100 \quad (A3)$$

In this equation,  $\overline{\omega_{HONO}}$  is the mean thermal HONO molecular velocity calculated by

$\overline{\omega_{HONO}} = \sqrt{8RT/\pi M}$ , where R, T, and M are the gas constant, absolute temperature and molecular weight. S/V is the surface-to-volume ratio ( $\text{cm}^2/\text{cm}^3$ ). The uptake coefficient  $\gamma_{HONO}$  was measured to be  $10^{-5}$  for soil surface and in the range of  $10^{-5}-10^{-3}$  for aerosol particle surface (Donaldson et al., 2014b; Wong et al., 2012). In addition,  $OH+HONO$  occurs at rates 1-2 orders of magnitude smaller than the uptake and therefore plays a minor role. The combined loss processes lead to a HONO lifetime of about 4 hours during the night.

HONO is generally assumed to be produced via heterogeneous  $NO_2$  hydrolysis disproportionation (R6) (Finlayson-Pitts et al., 2003), and the production rate of HONO is estimated by Eq. (A4), expressed in the unit of  $\text{ppbv-HONO ppbv}^{-1}\text{-NO}_2 \text{ s}^{-1}$ .

$$P_{HONO}^{night} = 0.5 \times R_{NO_2 \rightarrow HONO}^{surface} = 0.5 \times \gamma_{NO_2} \times \overline{\omega_{NO_2}} \times S/V \times 100 \quad (A4)$$

where  $\overline{\omega_{NO_2}}$  is the mean  $NO_2$  molecular velocity, S/V is the surface-to-volume ratio of particles, which could range from  $9.0 \times 10^{-6} \text{ cm}^2/\text{cm}^3$  to  $3.0 \times 10^{-4} \text{ cm}^2/\text{cm}^3$  for normally polluted areas and highly polluted areas respectively (Spataro and Ianniello,



2014). The S/V in biomass burning smoke plume has huge uncertainty; additionally, ground surface is also expected to play an important role in nighttime HONO production given our ground sampling location (Tuite et al., 2021; Scharko et al., 2017; Kebede et al., 2016; Stemmler et al., 2006), however, its S/V is not well defined/quantified.

505 Overall, considerable uncertainty remains regarding the rate coefficient of the heterogeneous processes in the daytime, as well as the HONO and NO<sub>2</sub> uptake coefficients and S/V ratio. This uncertainty, complicated further with large variability of fire behavior and emissions, make the HONO budget quantification extremely challenging.

## Appendix B. Quantification of isotopic fractionation factor

### B.1 Nighttime processes

#### 510 B.1.1 Isotopic fractionation of N and O associated with nighttime uptake

Surface uptake is the major sink for HONO during the night. Surface uptake of HONO has been found to be kinetically limited by bulk diffusion in particles containing viscous organic-water matrices, and incorporates two simultaneous processes: 1) reactive uptake of HONO on the bare particle/minerals surface, and 2) accommodation and reaction of HONO in the bulk aqueous layer, that is affected by pH and diffusion in the organic-water matrix (Donaldson et al., 2014b). The uptake coefficient of HONO is determined by the competition between these two processes as a function of fraction of water coverage on the surfaces  $\theta_{H_2O}$  ranging from 0 to 1 in Eq. (B1), where  $\gamma_0$  and  $\gamma_l$  are the reactive uptake coefficients of HONO onto particle (mineral/soil) surfaces at dry ( $\theta_{H_2O}=0$ ) and wet ( $\theta_{H_2O}=1$ ) conditions, respectively. Under completely dry conditions ( $\theta_{H_2O} = 0$  or relative humidity (RH) = 0%), the former process is dominant, and the isotopic fractionation can be estimated by the ratio of square root of inverse mass, which is caused by different thermal velocities ( $\omega_{HONO}$ ) of two isotopologues following Eq. (B2), where R is the gas constant, T is absolute temperature and M is the molecular weight. Thus, heavier isotopes are depleted in HONO, resulting in -10‰ and -20‰ for <sup>15</sup>E and <sup>18</sup>E respectively. By contrast, under wet conditions when RH is 30% which results in a monolayer water coverage on particle surfaces ( $\theta_{H_2O}=1$ ), the aqueous layer uptake becomes dominant and the wet uptake coefficient  $\gamma_l$  can be mechanistically simulated with a resistor model simplified as Eq. (B3) (Hanson, 1997; Pöschl et al., 2007). In Eq. (B3),  $\alpha$  is the accommodation coefficient describing the probability that a HONO molecule striking water-coated particle enters into the bulk liquid phase, and  $\Gamma_b$  is the solubility of HONO in the bulk water in the particles or soils.  $\Gamma_b$  can be calculated with Eq. (B4), where  $D_a$  is the apparent diffusion coefficient of HONO in the particle-water (soil(organiacs)-water) matrix, and  $\tau$  is the exposure time.  $H_{eff}$  is the effective Henry's law constant that depends on the absolute Henry's law constant for HONO, pH, and acid dissociation constants for HONO ( $K_{a1}$ ) and H<sub>2</sub>NO<sub>2</sub><sup>+</sup> ( $K_{a2}$ ).

$$530 \quad \gamma_{HONO} = (1 - \theta_{H_2O})\gamma_0 + \theta_{H_2O} \gamma_l \quad (B1)$$

$$\omega_{HONO} = \sqrt{\frac{8RT}{\pi M}} \quad (B2)$$

$$\gamma_l = \left(\frac{1}{\alpha} + \frac{1}{\Gamma_b}\right)^{-1} \quad (B3)$$

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$$\Gamma_b = \frac{4H_{eff}RT}{\omega_{HONO}} \sqrt{\frac{D_a}{\pi\tau}} \quad (B4)$$

$$k_{u-HONO} \propto \frac{\gamma_l \times \omega_{HONO}}{4} \quad (B5)$$

540 Taking the previously measured HONO  $\gamma_l$  of  $2 \times 10^{-5}$  as that for the light isotopologue, and  $\alpha$  of  $5.8 \times 10^{-5}$  as a constant (Donaldson et al., 2014b),  $\Gamma_b(\text{H}^{16}\text{O}^{14}\text{N}^{16}\text{O})$  is calculated to be  $1.36 \times 10^{-5}$  following Eq. (B3). As derived from Eqs. (B2) and (B4),  $\Gamma_b$  ratio between two isotopologues equals the ratio between the two molecular weights, and therefore  $\Gamma_b(\text{H}^{16}\text{O}^{15}\text{N}^{16}\text{O})$  and  $\Gamma_b(\text{H}^{16}\text{O}^{15}\text{N}^{16}\text{O})$  were calculated and used to derive the corresponding  $\gamma_l$  values. The fractionation factor associated with HONO uptake ( $\alpha_{u-HONO}$ ), defined as the ratio between heavy and light rate coefficients ( $k_H/k_L$ ), were calculated following the relationship determined by Eq. (B5). On the basis of this model, we estimate that the isotopic fractionation associated with the wet uptake process to be -2‰ and -4‰ for  $^{15}\text{E}$  and  $^{18}\text{E}$  respectively. From our calculation, RH clearly influences isotopic fractionation in the range of 0-30%, with wet uptake of HONO favouring a smaller kinetic isotope effect than dry uptake.

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### B.1.2 Isotopic fractionation of N and O associated with each nighttime HONO production pathway

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Heterogeneous conversion of  $\text{NO}_2$  to HONO has been widely accepted as the major secondary HONO production source during the night. However, the mechanism via which the conversion occurs remains disputed. Additionally, the kinetic isotopic fractionation factor (KIF) associated with this process has never been measured or calculated.  $\text{NO}_2$  hydrolysis (R6) on a variety of surfaces was determined to be a major source of HONO production. A compelling mechanism proposed by Finlayson-Pitts (Finlayson-Pitts et al., 2003) suggests R6 consists of a series of key steps including 1) dimer  $\text{N}_2\text{O}_4$  formation from recombination of two  $\text{NO}_2$  molecules in the gas phase, and uptake of gaseous  $\text{N}_2\text{O}_4$  by thin water film on the top surface layer, 2) aqueous phase isomerization of symmetric  $\text{N}_2\text{O}_4$  to asymmetric  $\text{ONONO}_2$  which is subsequently autoionizing to  $\text{NO}^+\text{NO}_3^-$  and reacting with  $\text{H}_2\text{O}$  to form HONO and  $\text{HNO}_3$ , and 3) desorption of HONO from aqueous to gas phase. Recently it was shown that reduction of  $\text{NO}_2$  on iron-bearing minerals and quinone-rich humic acid in soils and particulate matter (R7) leads to faster HONO production than  $\text{NO}_2$  hydrolysis. Although differing in reaction mechanism, the two possible pathways (R6 and R7) proceed in three steps including uptake of  $\text{NO}_2$  into surface aqueous layer, reactions in aqueous phase, and desorption of HONO from aqueous to gas phase. The first two steps are limited by aqueous diffusion, and it is reasonable to assume diffusion-limited processes in the aqueous phase create no KIF. As HONO desorption may involve hydrogen bond breaking

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of complex HONO•••(H<sub>2</sub>O)<sub>n</sub>, this process likely determines the KIF associated with the heterogeneous NO<sub>2</sub>-to-HONO conversion ( $\alpha_d$ ), as calculated by Eq. (B6), where  $\mu_l$  and  $\mu_h$  are the reduced mass for the light and heavy isotope containing pair, respectively (Shi et al., 2019). As a result,  $^{15}\epsilon$  and  $^{18}\epsilon$  are estimated to be -2.9‰ (n=1) to -4.5‰ (n=2) and -5.7‰ (n=1) to -8.9‰ (n=2) respectively. For the isotope mass balance modelling, mean values of  $^{15}\epsilon$  (3.7‰) and  $^{18}\epsilon$  (7.4‰) were generally used for steady-state R3, R6 and R7 under steady-state condition, and the low (n=2) and high (n=1) values were used to evaluate the lower and upper bound.

$$\alpha_d = \sqrt{\mu_l / \mu_h} \quad (\text{B6})$$

## B.2 Daytime

### B.2.1 HONO photolysis

The isotopic effect associated with photolysis (PIE) of HONO is calculated for the first time, following the  $\Delta$ ZPE-approach proposed by Yung and Miller (1997) to determine the PIE of N<sub>2</sub>O photolysis. In principle, the absorption spectrum for the same kind of electronic transition is expected to be similar in shape and intensity upon isotopic substitution, based on the assumption that the electronic potential energy surface is constant for each isotopologue. This assures the continuum levels (leading to photolysis) of the excited state are not significantly changed while the vibrational levels of the ground state vary with isotopic substitutions due to mass difference. The latter results in a lower ground state zero point energy (ZPE) for a heavy isotopologue than a light one, and cause blue shift in the absorption spectrum of the heavy isotopologue relative to the light one (Miller and Yung, 2000). When exposed to sunlight in the troposphere (>290 nm), HONO is known to feature a set of progressive absorption bands between 310-370 nm arising from electronic excitation  $\tilde{X}^1A' \rightarrow \tilde{A}^1A''$ , which result in HONO photolysis to OH and NO with nearly unity quantum yield (Cox et al., 1980; Suter and Huber, 1989). Under the aforementioned assumptions, we calculate the spectra blue shift of all three heavy isotopologues (HO<sup>15</sup>NO, H<sup>18</sup>ONO or HON<sup>18</sup>O) relative to that of HONO using the  $\Delta$ ZPE-approach as shown in Fig. B1 and Tables B1 and B2.

We calculated  $\Delta$ ZPE from  $1/2 \sum \Delta v_i$ , where  $\Delta v_i$  is the ground state vibrational frequency difference between the normal isotopologue (HONO) and the heavier isotopologue (HO<sup>15</sup>NO, H<sup>18</sup>ONO or HON<sup>18</sup>O) for each vibrational mode calculated via forced field by Monse et al. (MONSE et al., 1969). Note only HO<sup>15</sup>NO UV absorption was measured in previous study that reported an average blue shift of  $\sim 20 \text{ cm}^{-1}$  (8-40  $\text{cm}^{-1}$ ) relative to HONO, and this is consistent with our calculation (Table B1). Note trans-HONO/cis-HONO abundance ratio is 2.5 at room temperature (Suter and Huber, 1989), and the difference of  $\Delta$ ZPE for t-HONO and c-HONO are less than 0.5% for <sup>15</sup>N, and 2% for <sup>18</sup>O (Table B1). The effect of the difference on j calculation is negligible. With the measured absorption cross-section of HONO between 293-400 nm and the quantified blue shift of all three isotopologues, we calculate each photolysis rate coefficient following Eq. (B7), which is the integral of photolysis quantum yield  $\Phi_a(\lambda)$  ( $\approx 1$ ), absorption cross-section  $\sigma_a(\lambda)$  and solar actinic flux  $I(\lambda)$  as a function of wavelength.  $I(\lambda)$  is computed

with the radiation transfer model TUV (<http://www.acd.ucar.edu/TUV>, Madronich and Flocke, 1998)) at various locations and time during our sampling period. With these  $j$  values listed in Table B2 ( $j, j_{15N}, j_{18O1}, j_{18O2}$ ), the fractionation constant ( $^{15}\epsilon$  and  $^{18}\epsilon$ , ‰) associated with HONO photolysis is calculated following Eq. (B8), where  $j'$  and  $j$  are photolysis rate coefficient of heavy and light isotopologues respectively. Note we take the average of  $j_{18O1}$  and  $j_{18O2}$  as  $j_{18O}$  assuming the  $^{18}\text{O}$  is equally distributed between the two O-sites of HONO. Results shows  $^{15}\epsilon$  and  $^{18}\epsilon$  ranges from -1.9‰ to -4.3‰ (mean = -3.0‰,  $1\sigma = 0.7\%$ ,  $n = 18$ ) and -1.9‰ to -5.9‰ (mean = -3.1‰,  $1\sigma = 1.0\%$ ,  $n = 18$ ) respectively when HONO photolysis rate decreases from  $1.4 \times 10^{-3} \text{ s}^{-1}$  to  $5.3 \times 10^{-4} \text{ s}^{-1}$ .

$$j = \int \sigma_a(\lambda) \Phi_a(\lambda) I(\lambda) d(\lambda) \quad (\text{B7})$$

$$\epsilon = \left[ \left( \frac{j'}{j} \right) - 1 \right] \times 1000 \text{ ‰} \quad (\text{B8})$$

The negative values of  $^{15}\epsilon$  and  $^{18}\epsilon$  suggest both  $^{15}\text{N}$  and  $^{18}\text{O}$  will be enriched in the remaining HONO upon photolysis. Applying a Rayleigh fractionation model described by equation Eq. (B9), we obtain  $\delta^{15}\text{N}$  and  $\delta^{18}\text{O}$  of HONO ( $\delta_f$ ) as a function of the fraction of HONO left after photolysis ( $f$ ). The initial isotopic composition of HONO ( $\delta_0$ ) is taken from nighttime young smoke mean values in Table 1, as they are the best estimate of the fresh emission from the fires we investigated.

$$\ln(\delta_f + 1000\text{‰}) = \epsilon \ln(f) + \ln(\delta_0 + 1000\text{‰}) \quad (\text{B9})$$

## 615 **B.2.2 Isotopic fractionation of N and O associated with each daytime HONO production pathway**

OH + NO (R2) is a radical-radical recombination reaction, which is characteristic of stabilization of activated complex HONO\* via collisional energy transfer. This reaction type is characteristic of large KIF that enriches heavier isotopologues in the product at the low-pressure limit but almost no KIF at the high-pressure limit. The closer a reaction system is to the high-pressure limit, the less fractionation occurs (Chai and Dibble, 2014). Under the atmospheric pressure, the rate coefficient  $k_1$  is in the fall-off region but close to the high-pressure limiting rate coefficient  $k_{\text{atm}} = 1/3 k^\infty$  (Forster et al., 1995). Therefore, we expect a moderate positive  $^{15}\epsilon$  (~10‰) and  $^{18}\epsilon$  (~15‰) (Chai and Dibble, 2014; Burkholder et al., 2019). Kinetic isotopic fractionation (KIF) associated with photo-enhanced  $\text{NO}_2$  conversion is not known. Similar to the nighttime heterogeneous  $\text{NO}_2$  conversion, R3 is also expected to occur in the surface aqueous phase and the overall KIF is largely determined by that associated with the desorption of HONO from aqueous to gas phase. Thus,  $^{15}\epsilon_3$  and  $^{18}\epsilon_3$  are the same as that of R6 and R7 (Appendix B.1.2).

KIF associated with  $\text{HNO}_3/\text{p-NO}_3^-$  photolysis (R4) in the atmosphere has never been measured experimentally, and lack of  $\text{p-NO}_3^-$  absorption spectroscopy hinders calculation.  $^{15}\text{N}$  enrichment factor ( $^{15}\epsilon$ ) for photolysis of snow surface-adsorbed  $\text{HNO}_3$

630 under natural sunlight was theoretically determined to be  $\leq -47.9\%$  following the  $\Delta ZPE$  approach Yung and Miller (1997),  
 which well explained the  $^{15}\epsilon$  laboratory-measured for snow surface nitrate photolysis under the radiation of simulated sunlight  
 (Berhanu et al., 2014; Frey et al., 2009). If we take this  $^{15}\epsilon$  value, and the measured  $\delta^{15}\text{N}$  of nitrate (8‰ to 20‰), the HONO  
 produced from surface nitrate photolysis will be very negative (-38.9 to -27.5‰) within 2 hours of photolysis.  $^{18}\text{O}$  enrichment  
 factor ( $^{18}\epsilon$ ) for photolysis of snow surface-adsorbed  $\text{HNO}_3$  has been measured to range from 6.0‰ to 12.5‰ (Frey et al., 2009;  
 635 Berhanu et al., 2015).

### B.3 $\delta^{18}\text{O}$ transferring coefficient by different pathways

For  $^{18}\text{O}$ , in addition to KIF (enrichment factor,  $\epsilon_{\text{O}}^i$  in ‰),  $\delta^{18}\text{O}$  transferring from different reactants greatly influence  $\delta^{18}\text{O}$ -  
 640 HONO ( $\delta^{18}\text{O}_{\text{i,t}}$ ), especially when the two O atoms of HONO are derived from different reactants. That is, HONO formed from  
 different pathways (R2, R3, R6, R7) consists of  $\delta^{18}\text{O}$  of each O-containing reacting partner in proportion determined by  
 stoichiometry of reaction i, expressed with Eqs. (10)-(12). In R2, OH and NO equally contribute their O-atom to HONO  
 expressed with Eq. (10); In R3 and R7,  $\text{NO}_2$  is the exclusive O source of HONO while  $\text{H}_3\text{O}^+$  only contribute a  $\text{H}^+$  to HONO  
 (Ammann et al., 1998; George et al., 2005; Stemmler et al., 2006; Scharko et al., 2017; Kebede et al., 2016); In R6, the  
 645 hydrolysis mechanism discussed in Appendix B suggests the  $\text{H}_2\text{O}$ -derived  $\text{OH}^-$  and  $\text{NO}_2$ -derived  $\text{NO}^+\text{NO}_3^-$  equally contribute  
 their O-atom to HONO (Finlayson-Pitts et al., 2003).

$$\delta^{18}\text{O}_{2,\text{t}} = \frac{1}{2} \delta^{18}\text{O-OH} + \frac{1}{2} \delta^{18}\text{O-NO} \quad (\text{B10})$$

$$\delta^{18}\text{O}_{3(\text{or } 7),\text{t}} = \delta^{18}\text{O-NO}_2 \quad (\text{B11})$$

$$650 \quad \delta^{18}\text{O}_{6,\text{t}} = \frac{1}{2} \delta^{18}\text{O-H}_2\text{O} + \frac{1}{2} \delta^{18}\text{O-NO}_2 \quad (\text{B12})$$

During the day,  $\text{NO-NO}_2$  equilibrium is maintained via  $\text{NO}_2$  photolysis and NO oxidation by  $\text{O}_3$  and/or  $\text{RO}_2$  following R8 and  
 R9, and NO and  $\text{NO}_2$  are expected to possess similar  $\delta^{18}\text{O}$  and this can be expressed as  $\delta^{18}\text{O-NO}_x$ . During the night, due to  
 increased sink of  $\text{NO}_x$  and decreased  $\text{O}_3$  concentration,  $\delta^{18}\text{O-NO}_x$  is expected to be lower than during daytime.  $\text{NO}_x$  resulting  
 655 from R8 and R9 should carry  $\delta^{18}\text{O}$  of  $\text{RO}_2$  and  $\text{O}_3$  respectively via transfer; as  $\text{RO}_2$  and  $\text{O}_3$  have very different  $\delta^{18}\text{O}$  values  $\sim$   
 $+23\%$  and  $+117\%$  respectively, the competition between R8 and R9 critically affects  $\delta^{18}\text{O-NO}_x$ , as described by equations  
 Eqs. (11) and (12). OH radical in the troposphere has been calculated to be -35‰ depleted in  $^{18}\text{O}$  relative to  $\text{H}_2\text{O}$  as a result of  
 isotopic exchange at 298 K (Walters and Michalski, 2016); by taking the  $^{18}\text{O}$  values for summertime precipitation water in the  
 western US (-10‰ to -5‰) (Welker, 2000) and the  $\text{H}_2\text{O}$  liquid-to-vapor enrichment factor  $\epsilon_{g-l}$  of +9‰ at 298 K derived from  
 660 literature with Eq. (B13) (Michalski et al., 2012),  $\delta^{18}\text{O-OH}$  is estimated in the range of -35‰ to -30‰ if we ignore the unknown  
 KIF derived from OH oxidation reaction with the vast majority of atmospheric species. The overall  $\delta^{18}\text{O-HONO}$  is modelled  
 using the isotope mass balance model

$$\varepsilon_{g-l} = -7.68 + 6.71 \left( \frac{10^3}{T} \right) - 1.67 \left( \frac{10^6}{T^2} \right) + 0.35 \left( \frac{10^9}{T^3} \right) \quad (\text{B13})$$

665

*Data availability.* All data are available in the manuscript, the supplementary materials, or data repository (DOI: <https://doi.org/10.26300/k056-fs32>).

670 *Supplement.*

*Author contributions.* JD and MH conceived the research. JC, JD and MH designed the research. JC carried out field sampling, laboratory sample analyses, data analyses and figure production, as well as conceived and carried out the isotopic box modeling work. BA, CJ, WW, DB, EJ, JK, HM and EH contribute to field sampling. CB helped with laboratory sample analyses. JC wrote the paper. All authors contributed to the scientific discussions and preparation of the manuscript.

675 *Competing interests.* The authors declare that they have no conflict of interest.

*Disclaimer.* Any mention of brand names or manufacturers is for information purposes only and does not constitute an endorsement.

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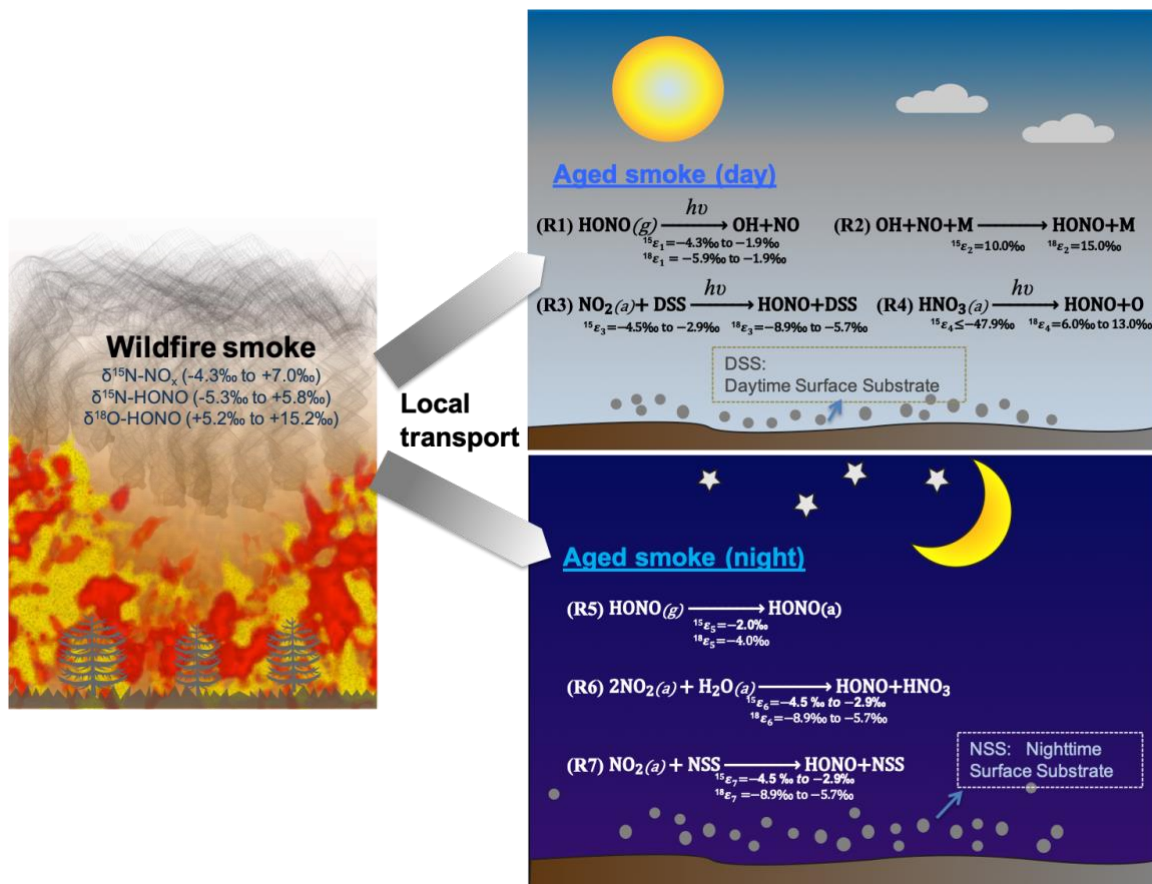
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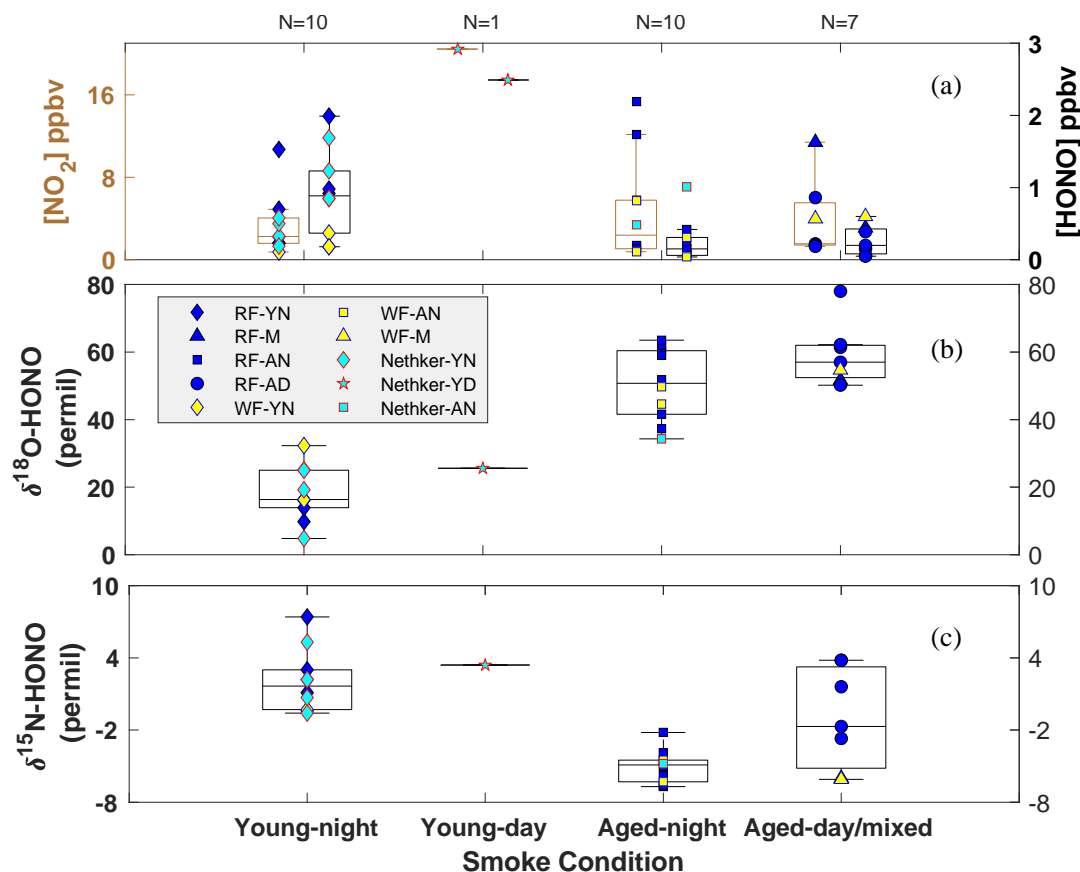


925 **Figure 1. The schematics of loss and secondary production of HONO in areas impacted by wildfire smoke in daytime (R1-R4) and nighttime (R5-R7).** We conducted our sample collection <30 km from the edge of the wildfires, with smoke ages ranging from a few minutes to half a day. M is bath gas including  $\text{N}_2$ ,  $\text{O}_2$ ,  $\text{CO}_2$ , etc. DSS is daytime substrate surfaces including terrestrial surfaces and aerosol particles that incorporate photoactive metal oxides (e.g.  $\text{TiO}_2$ ), humic like organics (e.g. quinone), etc. In essence, solar radiation induces reduction of these substrates with H, and this facilitates H abstraction

930 by  $\text{NO}_2$  (or H transfer). NSS is nighttime substrate surfaces (terrestrial and aerosol surfaces) containing iron-bearing minerals and/or humic acid (quinone). Note other sinks during both day and night (e.g.  $\text{OH} + \text{HONO}$ ) are negligible compared to the major sinks shown here. Isotopic enrichment factors for N and O result from kinetic isotopic effects associated with each

reaction and are calculated and expressed as  $^{15}\epsilon_i$  and  $^{18}\epsilon_i$ , where the subscript number  $i$  indicates the reaction number, and the superscripts 15 and 18 denotes the isotopic composition  $^{15}\text{N}/^{14}\text{N}$  and  $^{18}\text{O}/^{16}\text{O}$ , respectively.

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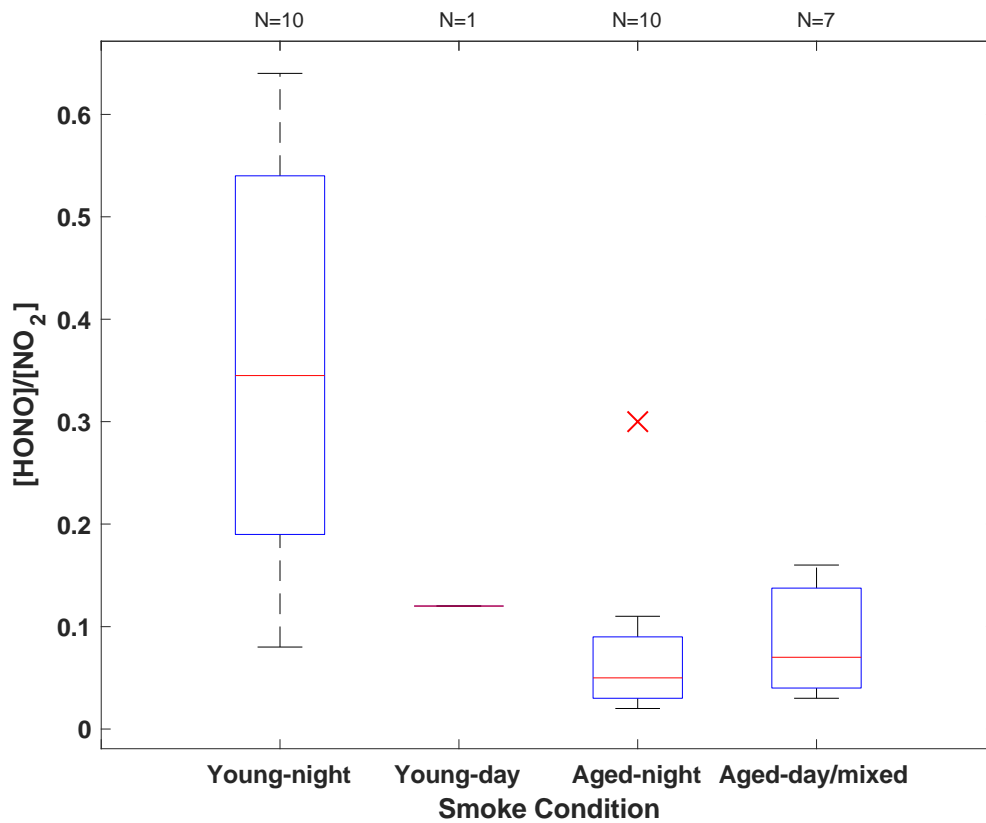


**Figure 2. Box-whisker plots for concentration of  $\text{NO}_2$  (left) and HONO (right) (a),  $\delta^{18}\text{O}$ -HONO (b),  $\delta^{15}\text{N}$ -HONO (c)**

**for each sample.** Individual data points are plotted within each box grouped by various field smoke conditions including young nighttime smoke (YN), young daytime smoke (YD), mixed daytime smoke (M) that contains smoke contributed by

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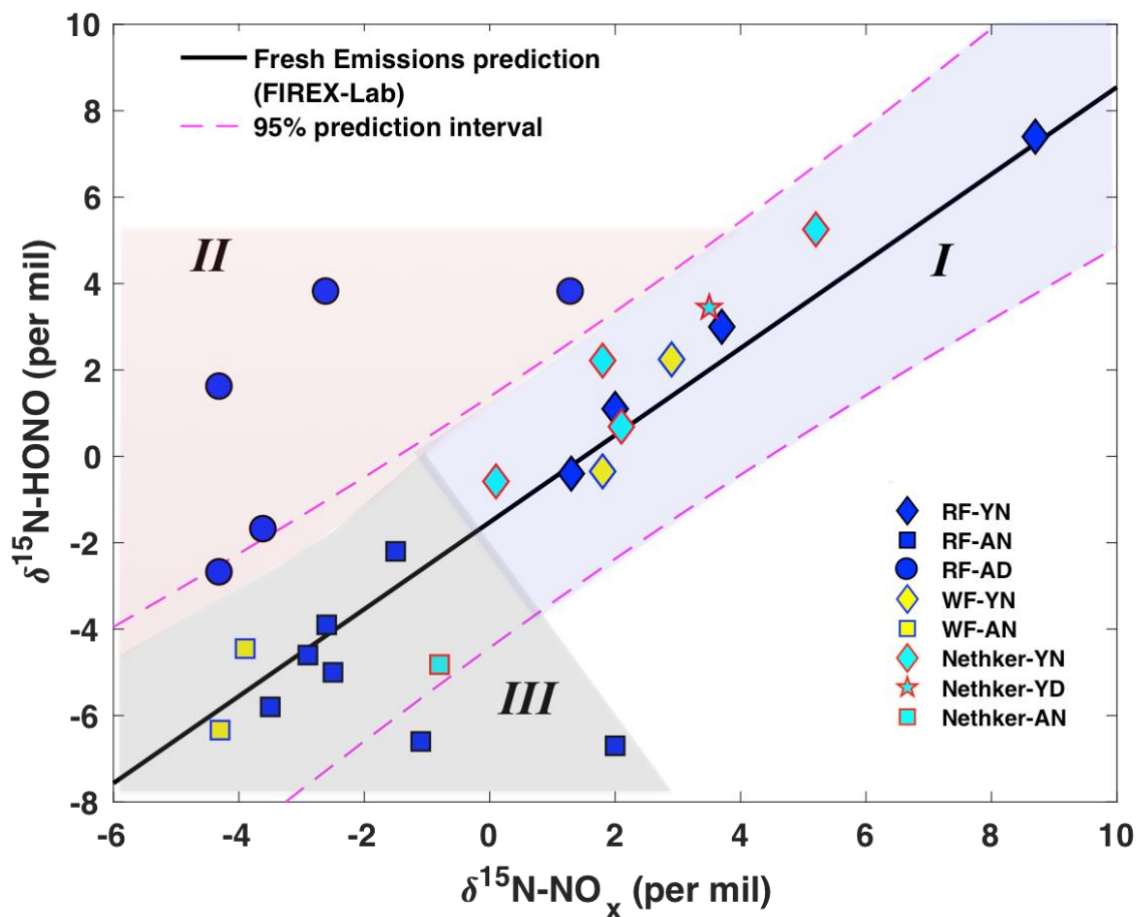
either night smoke or fresh smoke, aged nighttime smoke (AN), and aged daytime smoke (AD). Data from three wildfires are shown here, including Rabbit Foot (RF) fire during the 2018 WE-CAN campaign, Williams Flats (WF) fire and Nethker fire during the 2019 FIREX-AQ campaign.  $N$  is sample number measured for each condition. Each box-whisker presents the 5th, 25th, 50th, 75th, 95th percentile of sample values in each group.



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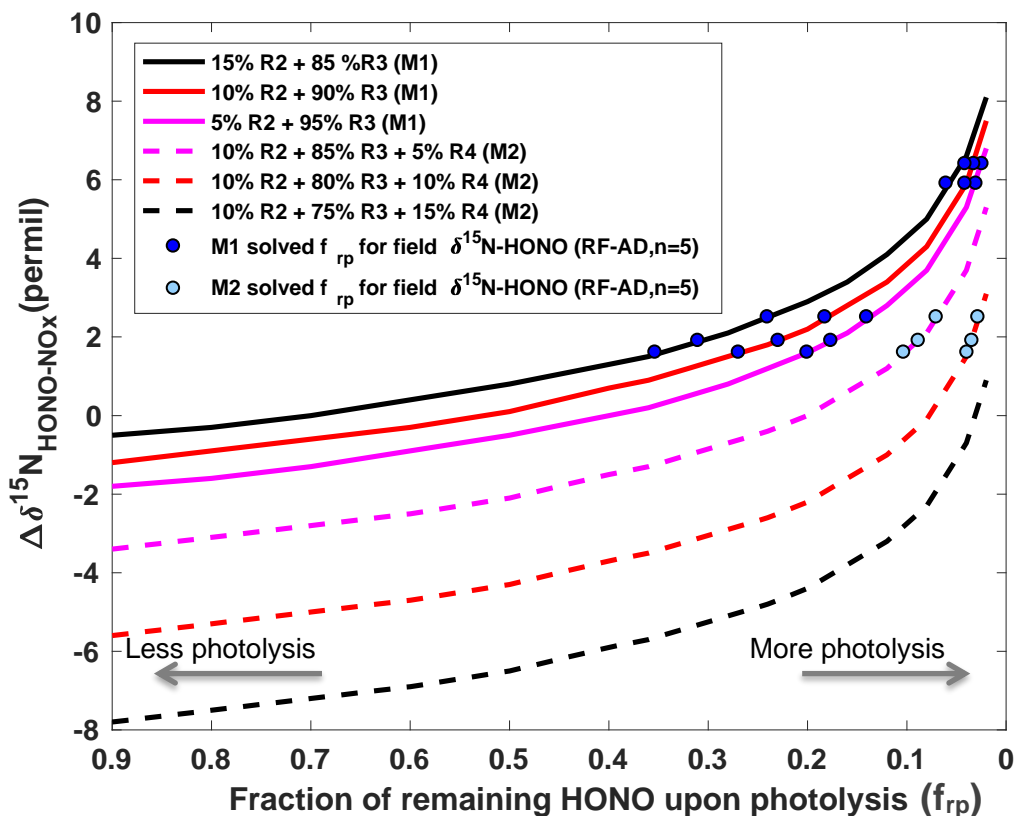
**Figure 3.** HONO/NO<sub>2</sub> concentration ratio summarized in box-whisker plot for each sampling condition. Red cross indicates an outlier. The whiskers from bottom to top represent 5%, 25%, 50%, 75% and 95% quartiles.





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**Figure 4. Relationship between wildfire-derived  $\delta^{15}\text{N-HONO}$  and  $\delta^{15}\text{N-NO}_x$ .** Samples from plumes of three wildfires including Rabbit Foot fire (RF; 2018), Williams Flats fire (WF; 2019) and Nethker fire (2019) are shown as different colors. Different symbols indicate different smoke conditions including young nighttime smoke (YN), young daytime smoke (YD), aged nighttime smoke (AN), and aged daytime smoke (AD). Note that the mixed smoke samples displayed in Fig. 2 are not shown here due to their large uncertainty. The black solid line ( $\delta^{15}\text{N-HONO} = 1.01\delta^{15}\text{N-NO}_x - 1.52$  ( $R^2 = 0.89$ ,  $p < 0.001$ )) is derived from lab-controlled burning emissions during the 2016 FIREX fire lab study (Chai et al., 2019), and within the 95% confidence interval (magenta dashed lines) predicts much of the field-based  $\delta^{15}\text{N-HONO}$  versus  $\delta^{15}\text{N-NO}_x$ . The field data are further grouped into three regimes—young smoke in both day and night (I, light purple shading), aged daytime smoke (II, pink shading), aged nighttime smoke (III, gray shading) based upon the  $\delta^{18}\text{O-HONO}$  results.



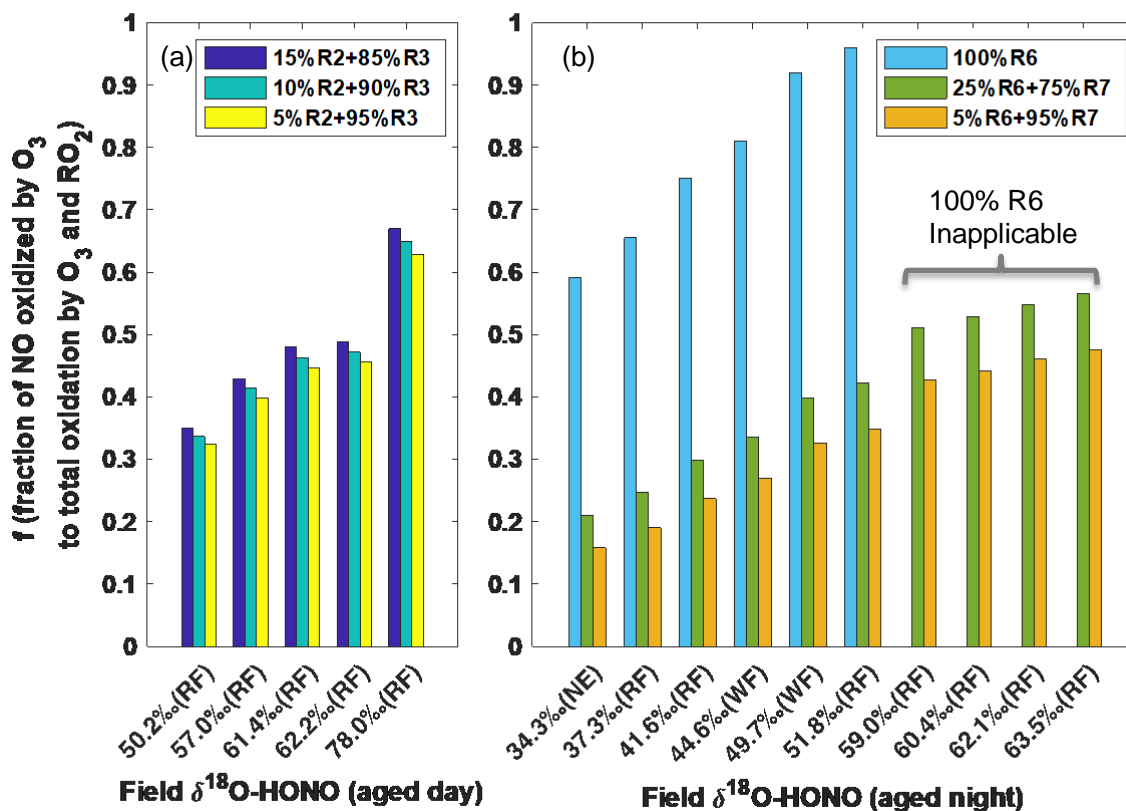
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**Figure 5. Modelling results of  $\delta^{15}\text{N}$  for aged daytime smoke via two plausible mechanisms (M1 and M2) for secondary**

**HONO production.** The isotope mass balance model (Eq. (1)) is used to simulate the  $\delta^{15}\text{N}$  difference ( $\Delta\delta^{15}\text{N}_{\text{HONO-NOx}} = \delta^{15}\text{N}_{\text{HONO}} - \delta^{15}\text{N}_{\text{NOx}}$ ) as a function of fraction of HONO remaining after photolysis ( $f_{\text{rp}}$ ) in a pseudo-photochemical steady

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state. The calculated kinetic fractionation factors used here are explained in Appendix B. In the first mechanism (M1, solid lines), R3 is the major HONO production pathway with varying relative contribution from R2 (see Fig. 1 for reactions), which is constrained as producing no more than 15% of the observed HONO concentrations. In the second mechanism (M2, dashed lines), nitrate photolysis (R4) is included in addition to R2 and R3 for HONO production. Taking the contribution of R2 of 10% as a constant, three scenarios were modelled by varying the relative contribution of R3 (75%-85%) and R4 (5%-15%).



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Figure 6. Model prediction of fraction of NO oxidized to NO<sub>2</sub> via O<sub>3</sub> to that via O<sub>3</sub> and RO<sub>2</sub> together ( $f_{O_3/(O_3+RO_2)}^{NO}$ ) on the basis of field-measured δ<sup>18</sup>O-HONO for aged daytime (a) and nighttime (b) smoke. During the day (a), the contribution of R2 to HONO production is varied from 5% to 15% following M1 in Fig. 5, and R3 accounts for the remaining secondary HONO contribution. The modelling results are shown in Fig. S5 and Table S3 in the supplement.

975 During the night (b), three scenarios with various contributions of R6 and R7 are modelled (Fig. S6 and Table S4).

$f_{O_3/(O_3+RO_2)}^{NO}$  is predicted to be over unity for the last four observed δ<sup>18</sup>O-HONO values if R6 is assumed as the only nighttime pathway.

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**Table 1.** Sampling condition and isotopic composition and concentration results for NO<sub>x</sub> and HONO for Rabbit Foot (RF) fire during the 2018 WE-CAN (a), as well Williams Flats (WF) fire and Nethker fire during the 2019 FIREX-AQ campaign (b). Smoke conditions include young nighttime smoke (YN), young daytime smoke (YD), mixed daytime smoke (MD), aged nighttime smoke (AN), and aged daytime smoke (AD). The conditions are determined primarily by comparing the field data with the lab data involving three factors:  $\delta^{18}\text{O}$ -HONO,  $\delta^{15}\text{N}$  relationship between HONO and NO<sub>x</sub>, HONO/NO<sub>x</sub> (or HONO/NO<sub>2</sub>) ratio, along with the smoke sampling locations. Specifically, significantly elevated  $\delta^{18}\text{O}$ -HONO indicates secondary production of HONO. Note during 2019 campaign, NO<sub>x</sub> concentrations were not measured due to instrumental breakdown.

(a)

Start time (MDT)	End time (MDT)	Fire (smoke condition)	$\delta^{18}\text{O}$ -HONO	$\delta^{15}\text{N}$ -HONO	$\delta^{15}\text{N}$ -NO <sub>x</sub>	[HONO] ppbv	[NO <sub>x</sub> ] ppbv	[NO <sub>2</sub> ] ppbv	HONO/NO <sub>2</sub>	HONO/NO <sub>x</sub>
8/9/18 10:11	8/9/18 19:10	RF (AD)	50.2‰	3.8‰	1.3‰	0.06	1.8	1.42	0.04	0.03
8/9/18 21:51	8/10/18 8:29	RF (AN)	37.3‰	-5.0‰	-2.5‰	0.06	1.3	1.35	0.04	0.05
8/10/18 9:50	8/10/18 20:26	RF (AD)	61.4‰	-1.7‰	-3.6‰	0.20	1.5	1.30	0.16	0.13
8/10/18 20:31	8/11/18 8:08	RF (AN)	60.4‰	-4.6‰	-2.9‰	0.15	1.4	1.39	0.11	0.11
8/11/18 22:43	8/12/18 9:38	RF (AN)	51.8‰	-5.8‰	-3.5‰	0.09	1.1	1.07	0.09	0.08
8/12/18 21:25	8/13/18 3:33	RF (AN)	62.1‰	-3.9‰	-2.6‰	0.26	0.8	0.81	0.33	0.33
8/13/18 3:53	8/13/18 7:05	RF (YN)	16.4‰	7.4‰	8.7‰	0.92	1.9	1.59	0.58	0.48
8/14/18 4:11	8/14/18 6:12	RF (YN)	16.1‰	-0.4‰	1.3‰	0.18	1.7	1.62	0.11	0.11
8/14/18 10:38	8/14/18 17:18	RF (AD)	57.0‰	1.6‰	-4.3‰	0.24	1.8	1.56	0.16	0.13

8/14/18	8/14/18	RF (AD)	78.0‰	3.8‰	-2.6‰	0.05	1.5	1.44	0.03	0.03
17:22	22:11									
8/15/18	8/15/18	RF (YN)	9.8‰	1.1‰	2.0‰	0.98	5.5	4.90	0.20	0.18
0:08	4:36									
8/15/18	8/15/18	RF (YN)	13.9‰	3.0‰	3.7‰	1.99	11.7	10.70	0.19	0.17
5:52	7:12									
8/15/18	8/16/18	RF (AN)	41.6‰	-2.2‰	-1.5‰	0.15	5.9	5.78	0.03	0.03
19:59	9:19									
8/16/18	8/16/18	RF (AD)	62.2‰	-2.7‰	-4.3‰	0.39	6.5	6.03	0.07	0.06
15:56	17:51									
8/16/18	8/17/18	RF (AN)	59.0‰	-6.7‰	2.0‰	0.42	15.6	15.34	0.03	0.03
21:22	6:25									
8/17/18	8/17/18	RF (M)	51.7‰	-6.0‰	-2.6‰	0.44	13.5	11.40	0.04	0.03
8:28	10:31									
8/17/18	8/18/18	RF (AN)	63.5‰	-6.6‰	-1.1‰	0.25	12.3	12.15	0.02	0.02
21:55	9:12									

995 (b)

Start time (MDT)	End time (MDT)	Fire (smoke condition)	$\delta^{18}\text{O}$ - HONO	$\delta^{15}\text{N}$ - HONO	$\delta^{15}\text{N}$ - NOx	[HONO] ppbv	[NO <sub>2</sub> ] ppbv	HONO/NO <sub>2</sub>
8/03/19 23:15:57	8/04/19 07:27:02	WF (AN)	44.6‰	-4.5‰	-3.9‰	0.31	5.7	0.05
8/04/19 18:25:49	8/05/19 09:40:08	WF (AN)	49.7‰	-6.3‰	-4.3‰	0.04	0.8	0.05
8/06/19 00:20:11	8/06/19 09:40:38	WF (YN)	16.3‰	-0.3‰	1.8‰	0.37	0.7	0.49
8/06/19 14:11:24	8/06/19 23:02:12	WF (M)	54.7‰	-6.1‰	-3.3‰	0.60	4.0	0.15
8/06/19 23:47:43	8/07/19 09:44:16	WF (YN)	32.3‰	2.2‰	2.9‰	0.18	2.2	0.08

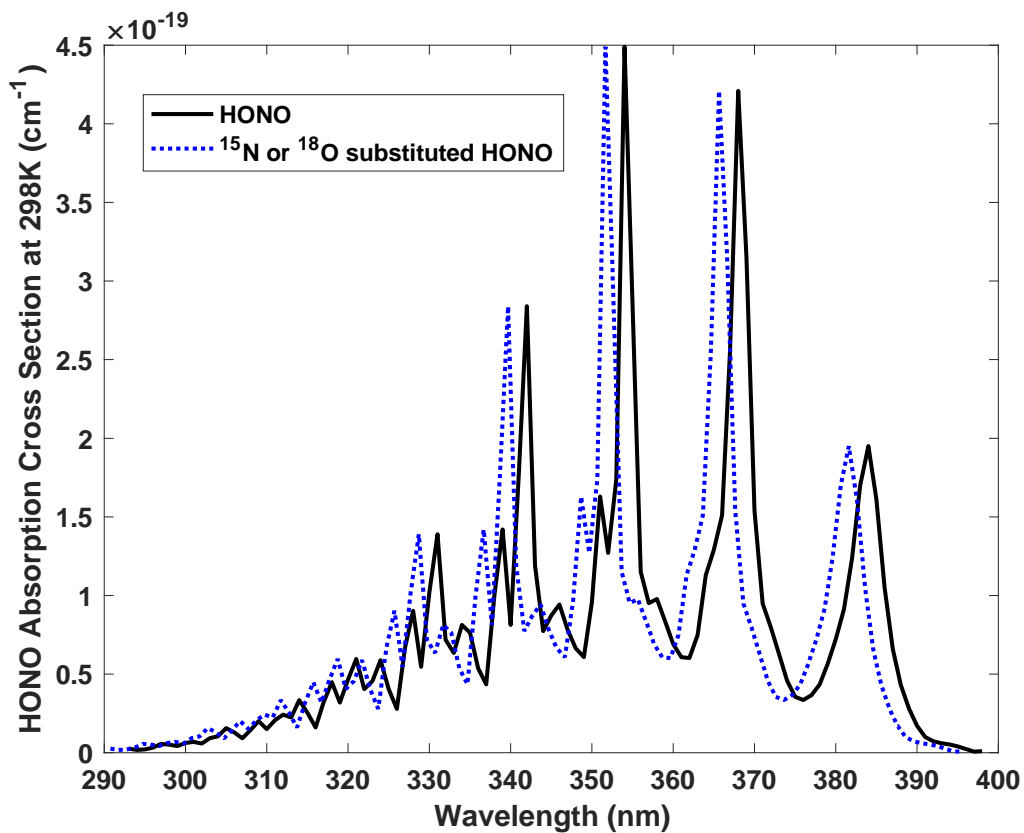
8/09/19	8/09/19								
12:32:42	14:56:34	Nethker(YD)	25.6‰	3.4‰	3.5‰	2.49	20.4	0.12	
8/10/19	8/11/19								
21:07:49	01:47:41	Nethker(YN)	25.1‰	2.2‰	1.8‰	1.23	2.3	0.54	
8/12/19	8/12/19								
03:24:15	11:24:47	Nethker(YN)	25.0‰	-0.6‰	0.1‰	1.69	3.5	0.48	
8/13/19	8/14/19								
21:38:03	01:28:27	Nethker(YN)	4.8‰	5.3‰	5.2‰	0.85	4.1	0.21	
8/15/19	8/15/19								
20:05:55	22:43:35	Nethker (AN)	34.3‰	-4.8‰	-0.8‰	1.01	3.4	0.30	
8/15/19	8/16/19								
22:57:07	06:28:04	Nethker(YN)	19.2‰	0.7‰	2.1‰	0.85	1.3	0.64	

**Table A1.** HONO budget estimation.

Start time (MDT)	Stop time (MDT)	[HONO] ppbv	[NO] ppbv	[NO <sub>2</sub> ] ppbv	j <sub>HONO</sub> s <sup>-1</sup>	[OH] = 1×10 <sup>6</sup> molecule cm <sup>-3</sup>		[OH] = 2×10 <sup>6</sup> molecule cm <sup>-3</sup>	
						k[OH]	P <sub>OH+NO</sub> /	k[OH]	P <sub>OH+NO</sub> /
						s <sup>-1</sup>	L <sub>HONO</sub>	s <sup>-1</sup>	L <sub>HONO</sub>
8/16/18	8/16/18								
15:56	17:51	0.39	0.51	6.03	1.2×10 <sup>-3</sup>	1.2×10 <sup>-5</sup>	0.01	2.4×10 <sup>-5</sup>	0.03
8/9/18	8/9/18								
15:38	19:10	0.06	0.41	1.42	1.1×10 <sup>-3</sup>	1.2×10 <sup>-5</sup>	0.08	2.4×10 <sup>-5</sup>	0.15
8/14/18	8/14/18								
10:38	17:18	0.24	0.22	1.56	1.4×10 <sup>-3</sup>	1.2×10 <sup>-5</sup>	0.01	2.4×10 <sup>-5</sup>	0.02
8/14/18	8/14/18								
17:22	20:11	0.05	0.07	1.44	6.1×10 <sup>-3</sup>	1.2×10 <sup>-5</sup>	0.03	2.4×10 <sup>-5</sup>	0.06
8/10/18	8/10/18								
9:50	20:26	0.20	0.22	1.30	1.1×10 <sup>-3</sup>	1.2×10 <sup>-5</sup>	0.01	2.4×10 <sup>-5</sup>	0.02

Note: L<sub>HONO</sub> = j<sub>HONO</sub>[HONO]

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**Figure B1.** Absorption wavelengths shift for  $\text{HO}^{15}\text{NO}$ ,  $\text{H}^{18}\text{ONO}$  and  $\text{HON}^{18}\text{O}$  compared with the most abundant form of HONO ( $\text{H}^{16}\text{O}^{14}\text{N}^{16}\text{O}$ ). The spectra of  $\text{HO}^{15}\text{NO}$ ,  $\text{trans-H}^{18}\text{ONO}$  and  $\text{trans-HON}^{18}\text{O}$  are blue shifted 0.23-0.43 nm, 0.21-0.39 nm, and 0.25-0.46 nm respectively spanning 293 to 398 nm. Note that the blue shift illustrated here is 2 nm (larger than the actual shift) in order to demonstrate the shift clearly.



**Table B1.** Vibrational frequencies of HONO and its isotopologues.

Vibrational mode	Vibrational frequency of trans-HONO (cm <sup>-1</sup> )			Vibrational frequency of cis-HONO (cm <sup>-1</sup> )		
	$\nu$ (t-HONO)	$\frac{\Delta\nu}{\nu}$ (t-HO <sup>15</sup> N) (t-H <sup>18</sup> ONO)	$\frac{\Delta\nu}{\nu}$ (t-HON <sup>18</sup> O)	$\nu$ (c-HONO)	$\frac{\Delta\nu}{\nu}$ (c-HO <sup>15</sup> N) (c-H <sup>18</sup> ONO)	$\frac{\Delta\nu}{\nu}$ (c-HON <sup>18</sup> O)
$\nu_1$ (O-H stretch)	3590.71	0.01	0.01	3426.2	0.01	0
$\nu_2$ (O=N stretch)	1699.76	32.5	39.28	1640.52	31.54	31.56
$\nu_3$ (HON bending)	1263.21	1.62	1.39	1302	0.57	6.24
$\nu_4$ (O-N stretch)	790.12	15.73	2.14	851.94	12.47	2.57
$\nu_5$ (O-N-O bending)	595.6	2.88	13.96	609	1.95	1.23
$\nu_6$ (torsion)	543	1.25	1.26	638.5	6.27	14.91
$\Delta ZPE$ (cm <sup>-1</sup> )		27.00	29.02		26.41	28.26

**Table B2.** Parameters used for TUV solar actinic flux modelling. The modelled enrichment coefficients for HONO photolysis for HO<sup>15</sup>NO, H<sup>18</sup>ONO and HON<sup>18</sup>O are presented in data repository (DOI: <https://doi.org/10.26300/k056-fs32>).

MDT	Latitude	Longitude	Altitude (m)	j(HONO) s <sup>-1</sup>	j(HO <sup>15</sup> NO) s <sup>-1</sup>	j(H <sup>18</sup> ONO) s <sup>-1</sup>	j(HON <sup>18</sup> O) s <sup>-1</sup>	ε <sup>15</sup> ‰	ε (H <sup>18</sup> ONO) ‰	ε (HON <sup>18</sup> O) ‰
8/16/18 3:56 PM	44.6726	-114.2339	1700	1.412E-03	1.408E-03	1.408E-03	1.407E-03	-2.6	-2.2	-3.0
8/16/18 5:51 PM	44.6726	-114.2339	1700	9.319E-04	9.287E-04	9.291E-04	9.283E-04	-3.4	-3.0	-3.8
8/9/18 3:38 PM	44.5048	-114.2320	1500	1.486E-03	1.482E-03	1.483E-03	1.482E-03	-2.5	-2.2	-2.9
8/9/18 7:10 PM	45.3870	-113.9619	1117	4.733E-04	4.724E-04	4.726E-04	4.722E-04	-1.9	-1.4	-2.3
8/14/18 10:38 AM	44.7173	-114.0226	1412	1.257E-03	1.253E-03	1.254E-03	1.253E-03	-2.8	-2.4	-3.3
8/14/18 5:18 PM	44.7173	-114.0226	1412	1.101E-03	1.098E-03	1.098E-03	1.097E-03	-3.1	-2.7	-3.5
8/14/18 12:00 PM	44.7173	-114.0226	1412	1.496E-03	1.492E-03	1.492E-03	1.491E-03	-2.5	-2.2	-2.9
8/14/18 1:30 PM	44.7173	-114.0226	1412	1.596E-03	1.592E-03	1.593E-03	1.592E-03	-2.4	-2.1	-2.8
8/14/18 3:00 PM	44.7173	-114.0226	1412	1.534E-03	1.530E-03	1.531E-03	1.530E-03	-2.5	-2.1	-2.8
8/14/18 5:22 PM	44.7173	-114.0226	1412	1.085E-03	1.082E-03	1.082E-03	1.081E-03	-3.2	-2.7	-3.4
8/14/18 8:11 PM	44.7173	-114.0226	1412	1.027E-04	1.023E-04	1.024E-04	1.023E-04	-4.2	-3.7	-4.7
8/14/18 6:40 PM	44.7173	-114.0226	1412	6.329E-04	6.304E-04	6.307E-04	6.302E-04	-3.9	-3.5	-4.3
8/10/18 9:50 AM	45.3870	-113.9619	1117	1.052E-03	1.049E-03	1.046E-03	1.045E-03	-3.2	-5.5	-6.3
8/10/18 8:26 PM	45.3870	-113.9619	1117	7.373E-05	7.342E-05	7.346E-05	7.339E-05	-4.3	-3.7	-4.7
8/10/18 12:00 PM	45.3870	-113.9619	1117	1.495E-03	1.491E-03	1.491E-03	1.490E-03	-2.6	-2.2	-2.9
8/10/18 2:00 PM	45.3870	-113.9619	1117	1.592E-03	1.588E-03	1.589E-03	1.587E-03	-2.4	-2.0	-2.8
8/10/18 4:00 PM	45.3870	-113.9619	1117	1.403E-03	1.400E-03	1.400E-03	1.399E-03	-2.7	-2.3	-3.0
8/10/18 6:00 PM	45.3870	-113.9619	1117	9.014E-04	8.984E-04	8.988E-04	8.980E-04	-3.4	-2.9	-3.8