

1 Author Response to Both Referee Comments:
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3 Response: We thank the two reviewers for thoughtful suggestions and constructive criticism that
4 have helped us improve our manuscript. Below we provide responses to reviewer concerns and
5 suggestions in blue font.
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7 **Reviewer 1:**

8 In the paper "Cloud Drop Number Concentrations over the Western North Atlantic Ocean:
9 Seasonal Cycle, Aerosol Interrelationships, and Other Influential Factors" by Dadashazar et al.,
10 the authors investigate cloud droplet number concentrations (Nd) and its influential factors on
11 multiple scales on the basis of diverse observational data sets. The analysis first describes high-
12 resolution aircraft measurements of a flight from the ACTIVATE campaign, and puts these high-
13 resolution measurements into a wider context with the description of the general seasonal cycles
14 of cloud/aerosol/meteorology of the region. The authors go into the details of aerosol size and
15 vertical distributions and calculate aerosol-cloud-interaction (ACI) statistics in different seasons.
16 Potential influential meteorological factors are described in two analyses, using a composite
17 approach by contrasting high and low Nd days, and by the application of a machine learning
18 algorithm. Some of the main findings of the paper are that ACI is generally strongest during DJF,
19 when Nd values tend to be highest, and that high Nd days are shown to feature systematically
20 different meteorology (e.g. stronger continental outflow) when compared to low Nd days.

21 The topic of the paper is highly relevant to the aerosol/cloud/climate community and of high
22 interest to the readership of ACP. The paper presents a thorough analysis of comprehensive
23 observational data sets that advances the scientific understanding of the observed Nd patterns of
24 the region. The paper is well written and structured and displays high-quality figures. I have only
25 some minor points the authors need to address and some specific remarks that the authors may
26 want to consider. I therefore recommend the manuscript for publication in ACP after minor
27 revisions.

28 **1 Minor Points**

- 29 1. I think the authors should discuss vertical velocity as one of the main drivers of Nd
30 variability in some more detail. The authors do this to some extent already in the
31 manuscript, but I think it is necessary to point out that this is likely an important factor
32 which is e.g. not provided as an input to the GBRT models. I would suggest that the
33 authors include near-surface wind speed as a proxy for boundary layer
34 turbulence/updrafts in the GBRT, especially since winds seem to be an important factor
35 in the composite analysis.
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37 Response: We have addressed this comment by including wind speed at 2 m (Wind_{2m}) as
38 an input parameter into the GBRT model. We also include near-surface wind direction
39 following the second reviewer's advice.
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41 The method section has been updated accordingly: "The following is the list of
42 thermodynamic/dynamic input parameters derived from MERRA-2: vertical pressure

43 velocity at 800 hPa (ω_{800}), planetary boundary layer height (PBLH), cold-air outbreak
 44 (CAO) index, wind speed and wind direction at 2 m ($wind_{2m}$ and $wind-dir_{2m}$), relative
 45 humidity (RH) in the PBL and free troposphere represented by RH_{950} and RH_{800} ,
 46 respectively. CAO index is defined as the difference between skin potential temperature
 47 (θ_{skt}) and air potential temperature at 850 hPa (θ_{850}) (Papritz et al., 2015). Updraft
 48 velocity plays a crucial role in the activation of aerosol into cloud droplets in warm
 49 clouds (Feingold, 2003; Reutter et al., 2009). Since the direct representation of updraft
 50 speed is not available from reanalysis data, near-surface wind speed (i.e., $wind_{2m}$) is used
 51 as a representative proxy parameter as an input parameter to the regression models.
 52 CERES-MODIS cloud parameters include liquid cloud fraction and cloud top height for
 53 low-level clouds. In addition, PERSIANN-CDR daily precipitation (Rain) was included
 54 as a relevant cloud parameter.”

55 Table 2 is also updated now:

56 **Table 2: List of input parameters used as predictor variables in the GBRT and linear models.**
 57 **Variables are grouped into three general categories.**

		Parameter
Aerosol		Sulfate surface mass concentration ($Sulfate_{sf-mass}$)
		Sea-salt surface mass concentration ($Sea-salt_{sf-mass}$)
		Dust surface mass concentration ($Dust_{sf-mass}$)
		Organic carbon surface mass concentration ($OC_{sf-mass}$)
Cloud		Low-level liquid cloud fraction ($CF_{low-liq.}$)
		Low-level liquid cloud-top effective height ($Cloud-top_{low-liq.}$)
		Precipitation rate (Rain)
Dynamic/ Thermodynamic		Cold-air outbreak index (CAO_{index}): $\theta_{skt} * -\theta_{850}$
		Relative humidity at 950 hPa (RH_{950})
		Relative humidity at 800 hPa (RH_{800})
		Vertical pressure velocity at 800 hPa (ω_{800})
		Wind speed at 2 m ($Wind_{2m}$)
		Wind direction at 2 m ($Wind-dir_{2m}$)
		Planetary boundary layer height (PBLH)

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59 *Skin potential temperature

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61 We also updated the regression results shown in Table 5 as follows:

62 “We show in Table 5 the performance of two linear models based on a single linear regression
 63 (with sulfate mass concentration), and a multi-regression that uses 14 input variables listed in Table
 64 2. In addition, Table 5 also lists the performance of the GBRT model that ingests 14 input variables,
 65 similar to the linear multi-regression model. The average R^2 scores of the test set for predicting Na
 66 based on a linear regression using only sulfate surface mass concentration were 0.17 and 0.09 in
 67 DJF and JJA, respectively. In contrast, R^2 between the multi-regression linear model and the test
 68 dataset increased to 0.28 and 0.25 for DJF and JJA, respectively. This increase in predictive

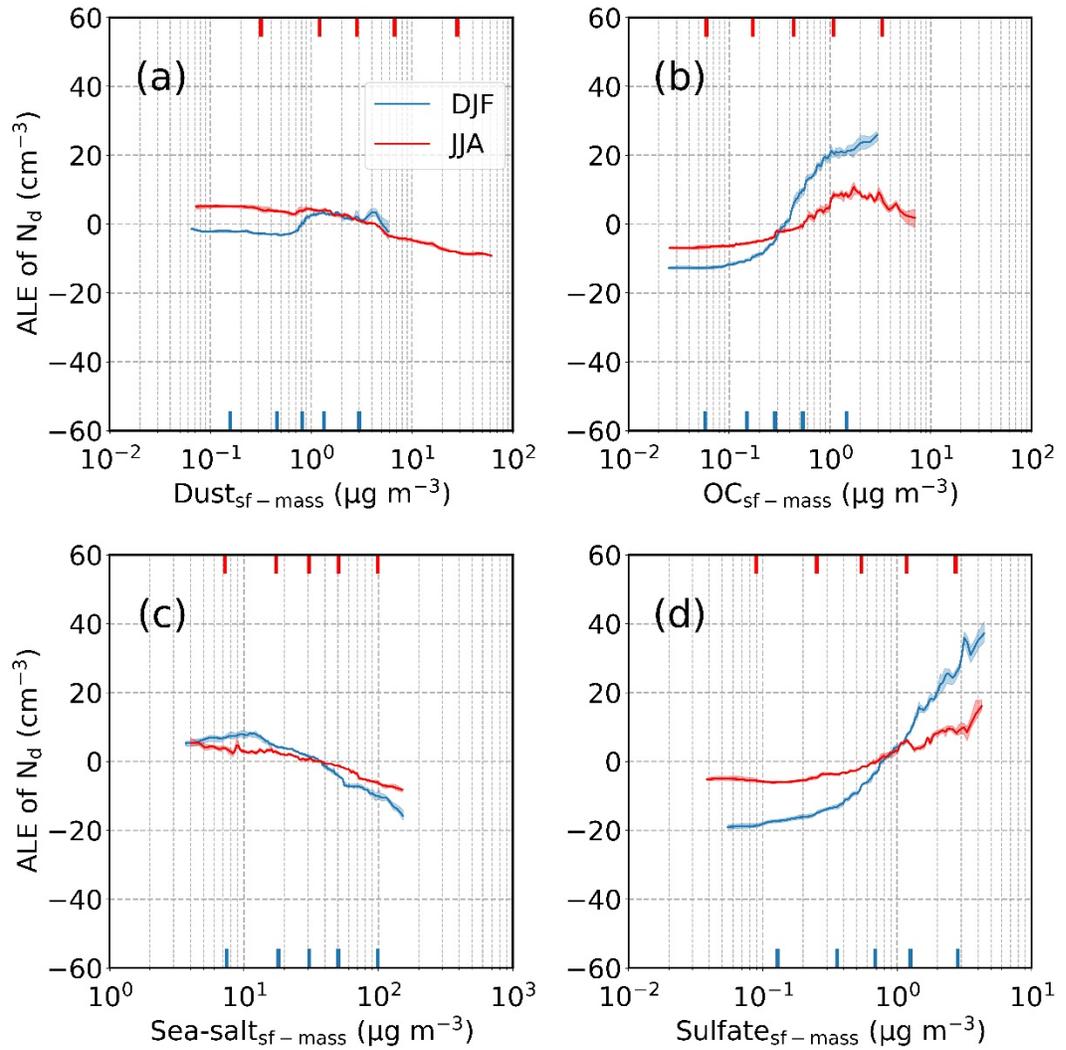
69 capability was helpful to reduce the gap between seasons by presumably accounting for factors
70 more important in JJA aside from surface concentration of sulfate. The R^2 scores increased even
71 more to 0.47 and 0.43 for DJF and JJA, respectively, for the GBRT model. Therefore, accounting
72 for non-linear relationships improved predictive capability in both seasons. It is important to note
73 that the GBRT model was robust in terms of overfitting and especially generalizability as R^2 values
74 of the test and validation sets were similar for both seasons.”

75 **Table 5: Performance of different models in predicting N_d assessed based on average**
76 **R^2 -scores on both validation and test sets. The models were fitted separately for DJF**
77 **and JJA seasons. Table 2 has the complete list of variables used in the GBRT model.**
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Model	Model type	Number of predictor variables	R^2 -score (DJF/JJA)	
			Validation set	Test set
$N_d \sim f(\text{Sulfate}_{\text{sf-mass}})$	Linear	1	0.17/0.09	0.17/0.09
$N_d \sim f(\text{Sulfate}_{\text{sf-mass}}, \text{CF}_{\text{low-liq.}, \dots})$	Linear	14	0.27/0.24	0.28/0.25
$N_d \sim f(\text{Sulfate}_{\text{sf-mass}}, \text{CF}_{\text{low-liq.}, \dots})$	GBRT	14	0.48/0.43	0.47/0.43

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82 ALE plots and associated discussion were also updated based on new results:

83 “The six parameters in Figure S21 (PBLH, RH_{950} , RH_{800} , Rain, $\text{Wind}_{2\text{m}}$, Wind-
84 $\text{dir}_{2\text{m}}$) did not reveal very pronounced trends with N_d in either season consistent with how
85 they did not rank highly in importance (Figure 13). Of particular interest is $\text{Wind}_{2\text{m}}$, which
86 is used here as a proxy variable for updraft speed in the marine boundary layer, which is
87 expected to have a high impact on N_d via its effect on in-cloud supersaturation. Although
88 the ALE plot of $\text{Wind}_{2\text{m}}$ suggested a small increase of about $\sim 10 \text{ cm}^{-3}$ in N_d as the wind
89 speed increased, $\text{Wind}_{2\text{m}}$ did not come out as a very important parameter in either seasons.
90 This may be due to the fact that environmental conditions representing updraft speed were
91 already included in parameters such as cloud fraction and CAO index. Another explanation
92 can be the shortcomings and high uncertainties associated with the use of $\text{Wind}_{2\text{m}}$ as a proxy
93 for updraft speed.”



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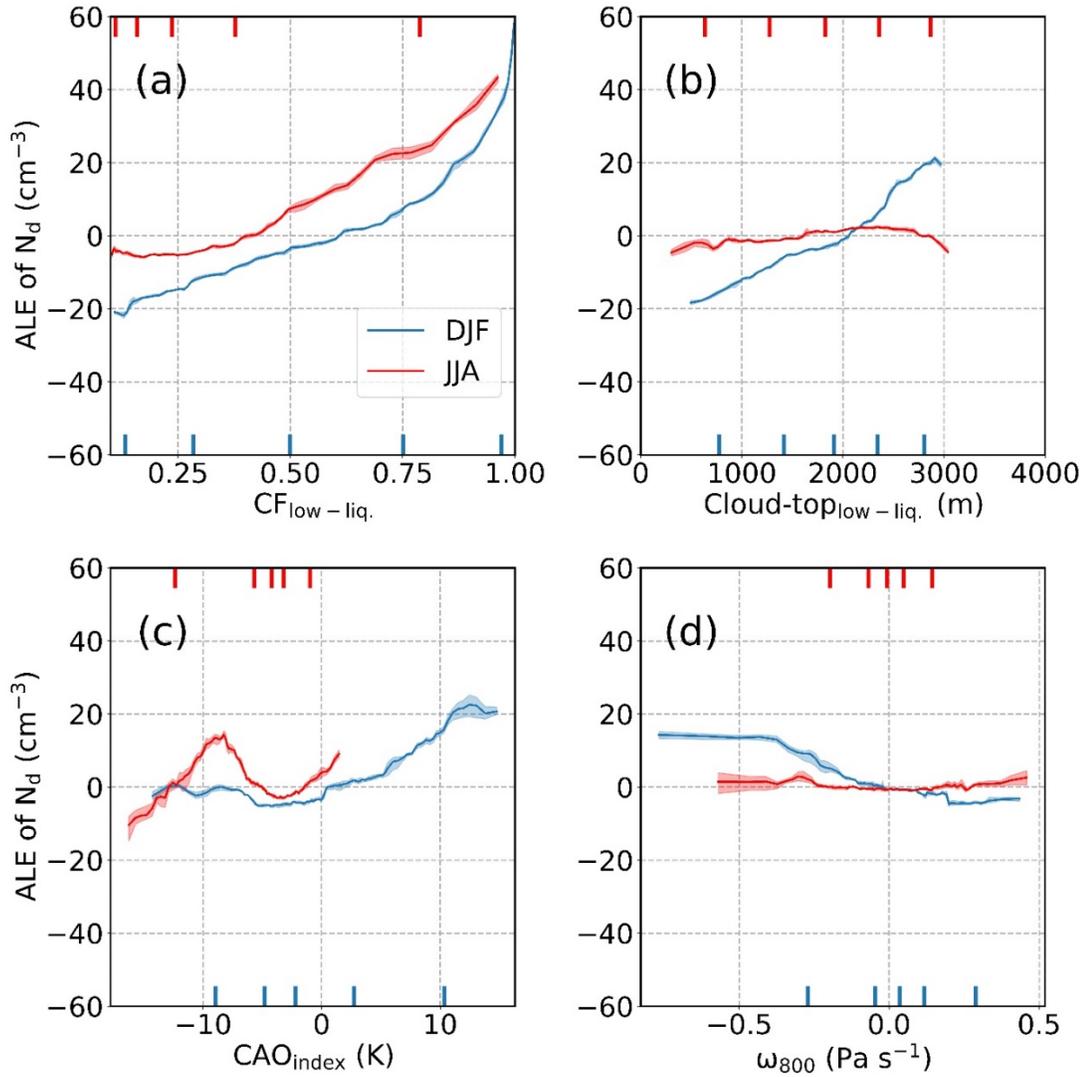
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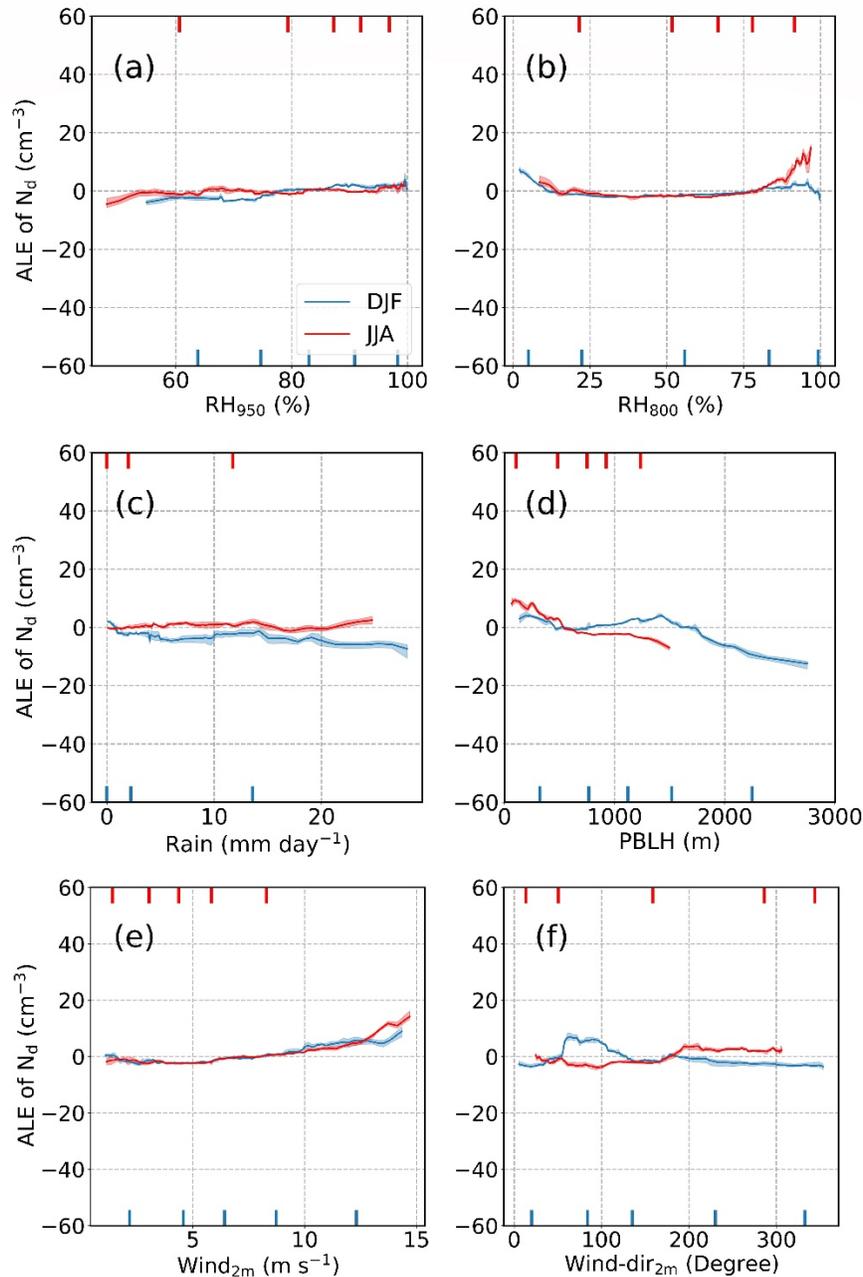
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Figure 15: Same as Figure 14 but for the following input parameters: (a) low-level liquid cloud fraction ($CF_{\text{low-liq}}$), (b) cloud-top effective height of low-level liquid cloud ($\text{cloud-top}_{\text{low-liq}}$), (c) cold-air outbreak (CAO) index, and (d) vertical pressure velocity at 800 hPa (ω_{800}).



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110 **Figure S21: Average local accumulated effect (ALE) profiles based on GBRT modeling of**
 111 **the following parameters: (a) relative humidity at 950 hPa (RH_{950}), (b) relative humidity at**
 112 **800 hPa (RH_{800}), (c) rain rate, (d) planetary boundary layer height (PBLH), (e) wind speed**
 113 **at 2 m (Wind_{2m}), and (f) wind direction at 2 m (wind-dir_{2m}). Blue and red profiles represent**
 114 **ALEs of DJF and JJA, respectively. Shaded areas show the ALE ranges stemming from the**
 115 **variability of the obtained models from the cross-validation resampling procedure. Markers**
 116 **on the bottom and top x-axes denote the values of 5th, 25th, 50th, 75th, and 95th percentiles for**
 117 **each input variable; note that the first three markers on the x-axes in panel (c) are very close**
 118 **and thus on top of each other.**

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2. In section 2.3 I am missing information on the hyperparameters of the GBRT models and how these are tuned during the training/validation phase. This is critical to be able to reproduce the results and informative for readers interested in the technical details of the model setup. In my opinion, this information could be provided for in a table and may be best suited in the supplement, though.

Response: We addressed this comment by adding a new table in SI file and also adding more detailed explanation in the method section (e.g., section 2.3) as follows: “Data were split into two sets: training/validation (70%) and testing (30%). Five-fold cross-validation was implemented to train the GBRT model using the training/validation data. Furthermore, both performance and generalizability of the trained models were tested via the aid of the test set, which was not used in the training process. Hyperparameters of the GBRT models were optimized through a combination of both random and grid search methods. Table S1 shows the list of important hyperparameters of the GBRT model and associated ranges tested via random and grid search methods. The optimized model hyperparameters can also be found in Table S1. The GBRT models were performed using the scikit-learn module designed in Python (Pedregosa et al., 2011).”

Table S1. Range of model hyperparameters tested during training/validation of the GBRT models through a combination of grid and random searches. Final model values are also listed in the last column.

Model parameter	Range of values tested	Final model values (DJF/JJA)
Learning rate	0.001-0.1	0.05/0.05
Number of estimators	100-5000	400/400
Maximum depth of a tree	2-35	9/11
Minimum number of samples to split an internal node	20-100	66/45
Minimum number of samples at a leaf node	20-60	31/66

3. What is the temporal relationship between the N_d and precipitation data? I believe it would be good to a) describe the time of satellite observations in subsection 2.2.1 and the precipitation data in 2.2.3 and b) discuss the potential influence of temporal offsets for the purpose of analyzing wet scavenging effects. I am also wondering why the authors did not chose to use information on precipitation from Cloudsat given they already use A-train data.

Response: We addressed this comment by adding the time of satellite observation in section 2.2.1: “Aqua observations used to estimate N_d were from the daytime overpasses of the satellite around 13:30 (local time).”

154 In section 2.2.3 we also described the temporal mismatch between N_d and precipitation
155 and pointed out the potential uncertainties: “It is important to note that we use daily
156 averaged PERSIANN-CDR precipitation and, therefore, there is some temporal mismatch
157 with the daily N_d value from MODIS-Aqua that comes at one time of the day. This can
158 contribute to some level of uncertainty for the discussions based on analyses involving
159 relationships between precipitation and N_d .”

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161 The lines below in section 4.1 were also updated:

162 “Furthermore, there was a general reduction in rain on low N_d days for most seasons except
163 SON, with rain enhancement on high N_d days except for DJF (Figure S6); this was
164 unexpected as wet removal was hypothesized to be a reason for reduced N_d for at least the
165 low N_d days. This may be attributed to the rain product being for surface precipitation (and
166 thus not capturing all drizzle) and for all cloud types, including more heavily precipitating
167 clouds deeper and higher than the low-level clouds examined for N_d . Another factor
168 potentially contributing to the observed counterintuitive trends is the temporal offset
169 between N_d estimations from MODIS-Aqua and precipitation data from PERSIANN-
170 CDR.”

171 The reviewer raised a good point about why we did not use precipitation data from
172 Cloudsat, which is part of A-train constellation and naturally temporally synchronized
173 with MODIS observations. We had reasons for this choice, with a major one being that
174 Cloudsat does not afford the spatial coverage we desired for our analyses. We had greater
175 flexibility using PERSIANN, however, we note that we are interested in future work to
176 do more detailed types of analyses that could be more catered to the strengths of Cloudsat
177 that are not available from other rain datasets.

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179 4. I think it would be beneficial to briefly comment if the vertical distribution of aerosols at
180 nighttime (used here) is expected to be significantly different from the daytime (rest of
181 the data sets analyzed here).

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183 Response: We added some comments regarding the use of nighttime CALIOP
184 observations in the beginning of section 3.4: “Vertical profiles of aerosol extinction
185 coefficient estimated from CALIOP nighttime observations are shown in Figure 4 for the
186 six sub-domains. Shown also are the seasonally representative planetary boundary layer
187 heights (PBLHs) from MERRA-2, with numerical values of both PBLH and fractional
188 AOD contributions to the PBL and FT in Table 3. Although here we used nighttime
189 observations from CALIOP because of having higher signal to noise ratio than daytime
190 observations, we expect the general seasonal trends discussed here to remain the same
191 regardless of the observation time.”

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193 For other work related to another project, we compared extensive CALIOP data for
194 daytime and nighttime conditions for the same region and saw no major qualitative
195 differences.

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5. This is just to initiate discussion: In the air-quality community, it is well known that the seasonal cycle of satellite-AOD and near-surface particle concentrations over continental regions frequently show contrasting seasonal cycles (e.g. Koelemeijer et al. 2006, 10.1016/j.atmosenv.2006.04.044). This is caused by effects of PBLH and BL humidity (Stirnberg et al.2018, 10.3390/rs10091353), and can be corrected for to some degree (Arvani et al. 2016, 10.1016/j.atmosenv.2016.06.037). I believe that at least qualitatively there is something to learn from these findings that have implications for the ACI community (and this paper specifically) as well, especially in studies covering continental regions or regions of strong continental outflow. I don't think the authors have to discuss this in their paper, but it may be a useful discussion to have in the ACI community and in my opinion links well to the findings presented here. The authors may chose ignore, comment or discuss this point as they find most useful.

Response: We thank the reviewer for this idea and important point. We aim to consider this line of discussion and thought for potential future lines of work and do not make changes to the current manuscript to address this issue.

2 Specific Remarks

1. 58 In my opinion, the word "potentially" does not apply to "enhanced cloud albedo" in the case of increased Nd and constant LWP, but only applies to the latter two points.

Response: Fixed as follows: "It is widely accepted that warm clouds influenced by higher number concentrations of aerosol particles have elevated Nd and smaller drops (all else held fixed), resulting in enhanced cloud albedo at fixed liquid water path (Twomey, 1977), and potentially suppressed precipitation (Albrecht, 1989) and increased vulnerability to overlying air resulting from enhanced cloud top entrainment (Ackerman et al., 2004)."

1. 206 There is a typo: "supermicromemter"

Response: Fixed : "Estimation of supermicrometer particles from FCDP measurements was performed after conducting the following additional screening steps to minimize cloud droplet artifacts:"

1. 209 Please use SI units (760 torrs)

Response: Fixed: "data collected during ACB and BCT legs were excluded. CCN, LAS, CPC, and AMS measurements are reported at standard temperature and pressure (i.e., 273 K and 101.325 kPa) while FCDP and 2DS measurements correspond to ambient conditions."

234 1. 385 I think Stier (2016, 10.5194/acp-16-6595-2016) is a relevant source that should be
235 cited here.

236 Response: Thanks for the great suggestions. It is added now: “While previous studies
237 have pointed to the limitations of AOD as an aerosol proxy (e.g. Stier, 2016; Gryspeerd
238 et al., 2017; Painemal et al., 2020), ...”
239

240 1. 482 ”usually always” - please remove one of them.

241 Response: Fixed: “Table 4 shows that DJF always exhibits the highest ACI values
242 regardless of the aerosol proxy used, consistent with a stronger aerosol indirect effect in
243 DJF over East Asia.”

244
245 **Reviewer 2:**

246 This paper attempts to explain the seasonality of cloud drop number concentrations (N_d) off the
247 coast of the USA and Canada with a focus on explaining why N_d is highest in the DJF season,
248 but the aerosol optical depth (AOD) is lowest in that season. The paper presents some useful
249 analysis and I think it should be published once the concerns have been addressed. However,
250 there are places where the results are not fully explained, some places where there should be a
251 more quantitative analysis and some results that are in the supplementary that should be in the
252 main text. Some of the arguments could also be made more clearly. It seems that aerosol is more
253 efficient at making cloud droplets in DJF, seemingly because of the prevalence of the trade
254 cumulus and stratocumulus conditions (with high low-altitude cloud fractions). However, the
255 paper doesn't quite get the bottom of why this should be – perhaps more discussion on this is
256 warranted although it may be case that we don't quite know yet. Here are some suggestions that
257 might help to get closer to an answer :

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259 • It would be interesting to examine what the important predictors are in the subset of data
260 with high low altitude cloud fractions. E.g., do the aerosol parameters then become
261 equally important between the seasons once we are in the cumulus/stratocumulus regime
262 (particularly CCN and sulphate surface mass)? Is N_d similar in DJF and JJA for the high
263 cloud fraction subset indicating that it is mainly the prevalence of the low cloud
264 conditions in DJF that cause the N_d difference? You could also do a similar analysis for
265 the subset with small low altitude cloud fractions (more typical of JJA perhaps).
266

267 Response: Thank you for the great suggestion. We ran the GBRT models for subsets of
268 data with varying cloud fractions ($CF_{low-liq}$: 0.2-0.4 and ≥ 0.7). The results suggested that
269 in high cloud coverage conditions (i.e., $CF_{low-liq} \geq 0.7$), sulfate surface mass
270 concentrations were the most important factor regardless of season with N_d showing very
271 similar relationships (and sensitivity) to sulfate surface concentrations in both seasons. In
272 contrast, different aerosol parameters appeared as being more important parameters in
273 DJF and JJA when only data with cloud fractions between 0.2 and 0.4 were included in
274 the GBRT model. In JJA, organic carbon was the most important factor while sulfate was
275 the most important parameter in DJF. It should be noted that for the low cloud fraction
276 model run, N_d exhibited less sensitivity to sulfate in JJA than DJF, which is similar to the

277 results of the original model run including all data points with cloud fraction greater than
278 0.1.

279 We added the results of these two sensitivity tests in the SI file and we updated
280 the text in the section 4.2 as follows:

281 “The results of regression analysis highlight the high sensitivity of N_d to cloud fraction
282 regardless of season. As discussed earlier, this can be attributed largely to two factors: (i)
283 the relationship between cloud type (e.g., stratocumulus, shallow cumulus) and cloud
284 fraction, which can, in turn, influence cloud microphysical properties like N_d ; and (ii)
285 uncertainties associated with N_d estimates from satellite observations that can result in
286 negative biases in N_d for low cloud coverage conditions. To further test the relative
287 influence of various variables at different cloud fractions, two sensitivity tests with GBRT
288 modeling were conducted using subsets of data with varying cloud fraction ($0.2 \leq CF_{\text{low-liqu.}} \leq 0.4$
289 and $CF_{\text{low-liqu.}} \geq 0.7$).

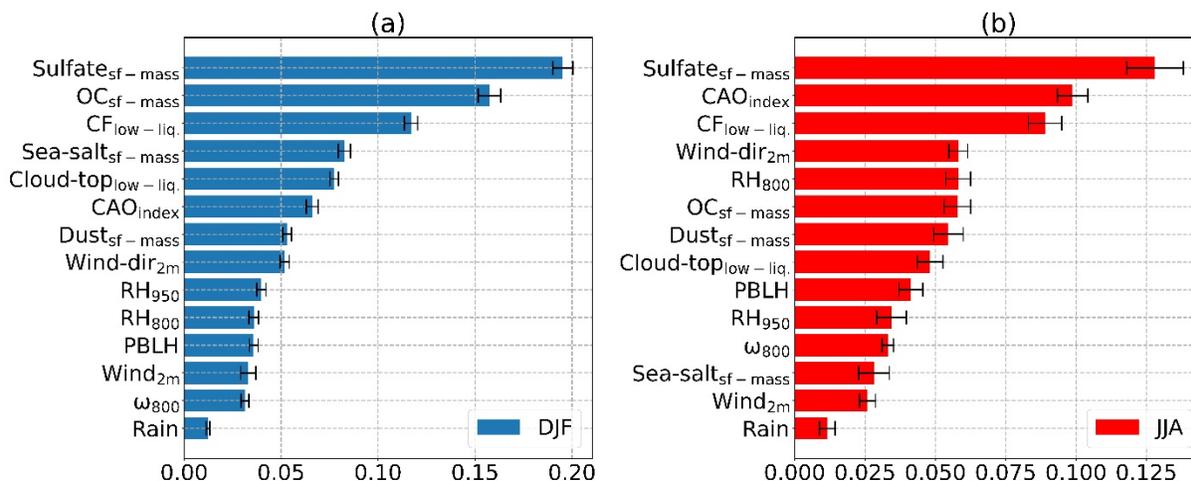
290 Beginning with results for $CF_{\text{low-liqu.}} \geq 0.7$ (Figures S22-25), the average R^2 -scores for
291 validation and test sets for these runs were 0.47/0.39 (DJF/JJA) and 0.49/0.38 (DJF/JJA),
292 respectively. A feature that stands out is that for both DJF and JJA, surface mass
293 concentrations of sulfate became the most important factor. ALE plots presented in Fig.
294 S23 also suggested that N_d has a very similar sensitivity to sulfate concentration in high
295 cloud coverage conditions regardless of season in contrast to the results of the original run
296 where N_d was more sensitive to the changes in sulfate level in DJF than JJA. These results
297 are in agreement with previous studies where N_d values for marine boundary layer clouds
298 were highly sensitive to sulfate concentrations at the level close to cloud base (Boucher
299 and Lohmann, 1995; Lowenthal et al., 2004; Storelvmo et al., 2009; McCoy et al., 2017;
300 McCoy et al., 2018; MacDonald et al., 2020). The second most important factor for DJF
301 was the surface mass concentrations of organic carbon followed by $CF_{\text{low-liqu.}}$ and sea-salt
302 surface mass concentrations. On the other hand, the second most important factor in JJA
303 was CAO index followed by $CF_{\text{low-liqu.}}$ and wind direction. ALE plots presented in Figs.
304 S23-25 showed similar relationships between N_d and input parameters as observed for the
305 original runs where full datasets were used as the input.

306 Figure S26 shows the results of the GBRT model using input data with cloud fraction
307 between 0.2 and 0.4, the condition relatively more prevalent in JJA than DJF. The average
308 R^2 -scores for validation and test sets for these runs were 0.30/0.30 (DJF/JJA) and 0.33/0.31
309 (DJF/JJA), respectively. It is interesting to see that for both seasons an aerosol parameter
310 emerged as the most important factor. Mass concentrations of OC appeared as the most
311 important factor in JJA (the fourth most important factor in DJF) while in DJF, sulfate
312 concentration came out as the most important factor (the fourth most important factor in
313 JJA) consistent with the results of previously discussed models for DJF. It should be noted
314 that ALE plots also suggested less sensitivity of N_d to sulfate in JJA than DJF, similar to
315 the results observed in the original model run including all the data points. The second
316 most important factor in DJF turned out to be the cloud-top effective height of low-level
317 liquid clouds followed by CAO index. On the other hand, CAO index was the second most

318 important factor in JJA followed by PBLH. ALE plots presented in Figs. S27-29 also
319 showed similar qualitative trends observed in original and high cloud coverage runs.”

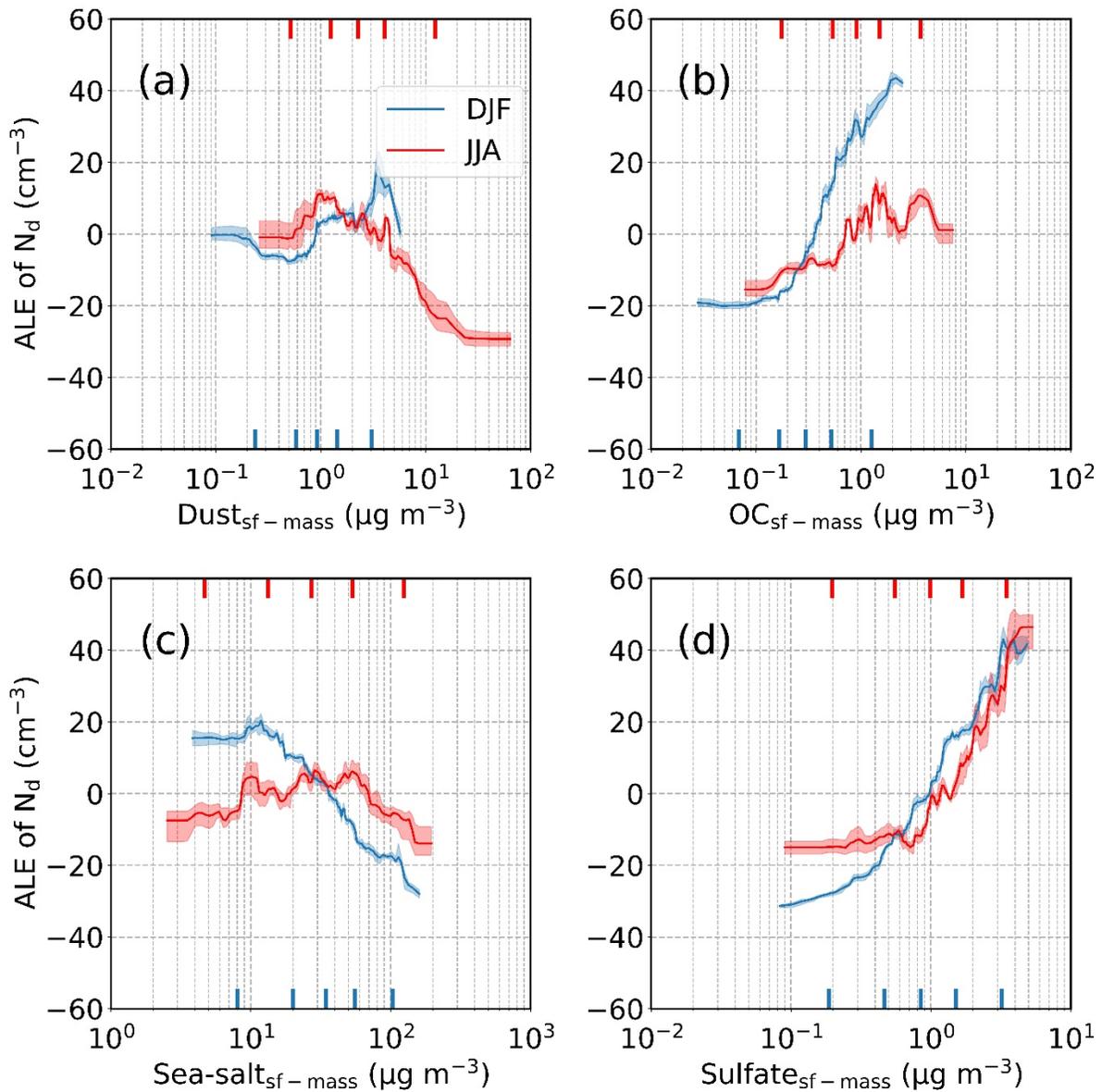
320 Here are the results for these new runs added to SI file:

321 “



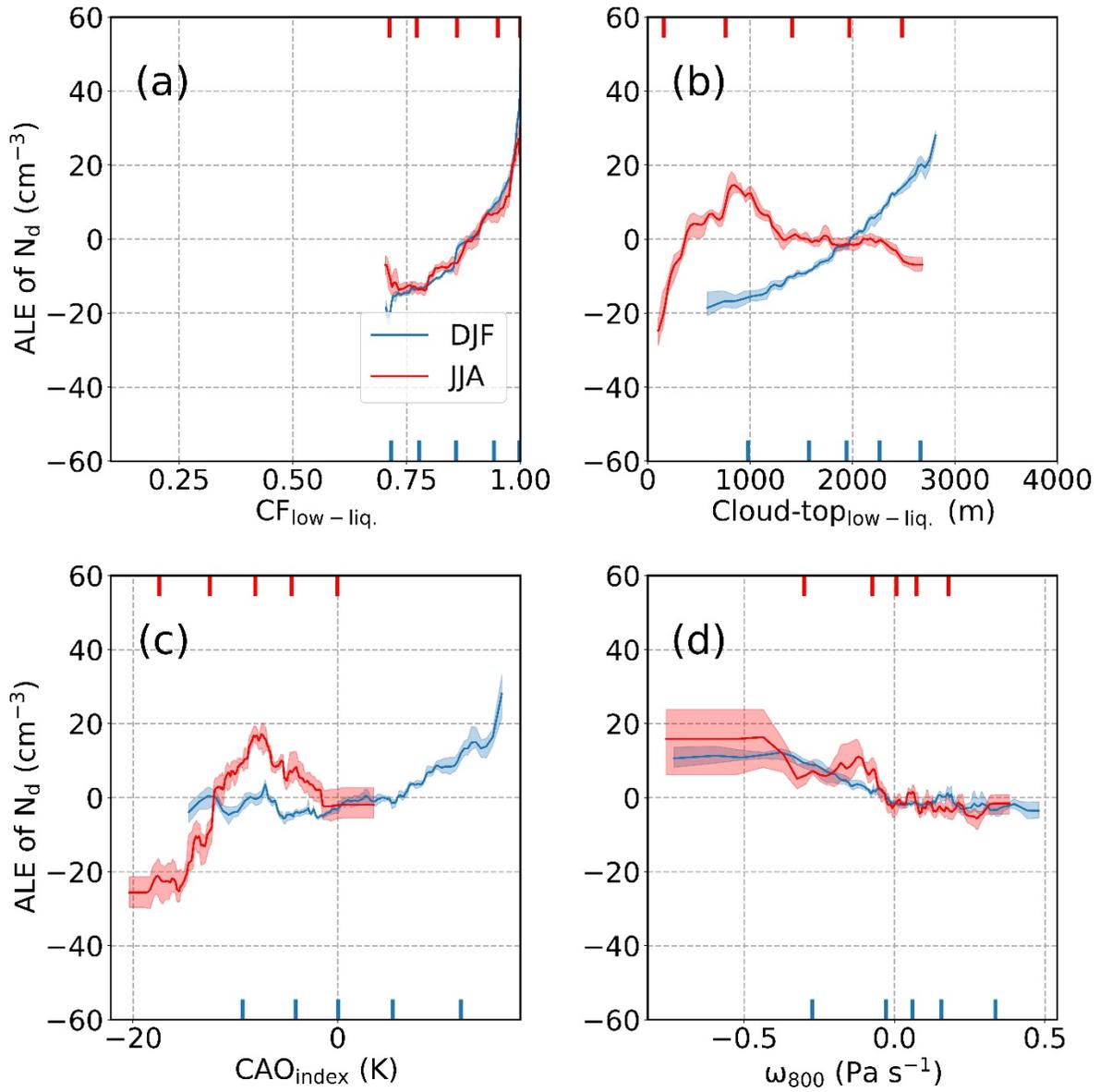
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323 **Figure S22: Average permutation feature importance of input parameters for (a) DJF and**
324 **(b) JJA based on GBRT models trained in each season on subsets of data including only**
325 **samples with low-level liquid cloud fraction greater than or equal to 0.7 (i.e., $CF_{low-liq.} \geq 0.7$).**
326 **Feature importance values were calculated based on using the test set. Error bars exhibit the**
327 **range of feature importance values stemming from the variability of the obtained models**
328 **from the cross-validation resampling procedure.**

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 331 **Figure S23: Average local accumulated effect (ALE) profiles based on GBRT modeling for**
 332 **surface mass concentrations of the following parameters: (a) dust, (b) organic carbon, (c)**
 333 **sea-salt, and (d) sulfate. ALE profiles were based on GBRT modeling on subsets of data**
 334 **including only samples with low-level liquid cloud fraction greater than or equal to 0.7 (i.e.,**
 335 **$CF_{\text{low-liq.}} \geq 0.7$). Blue and red profiles represent ALEs of DJF and JJA, respectively. Shaded**
 336 **areas show the ALE ranges stemming from the variability of the obtained models from the**
 337 **cross-validation resampling procedure. Markers on the bottom and top x-axes denote the**
 338 **values of 5th, 25th, 50th, 75th, and 95th percentiles for each input variable.**

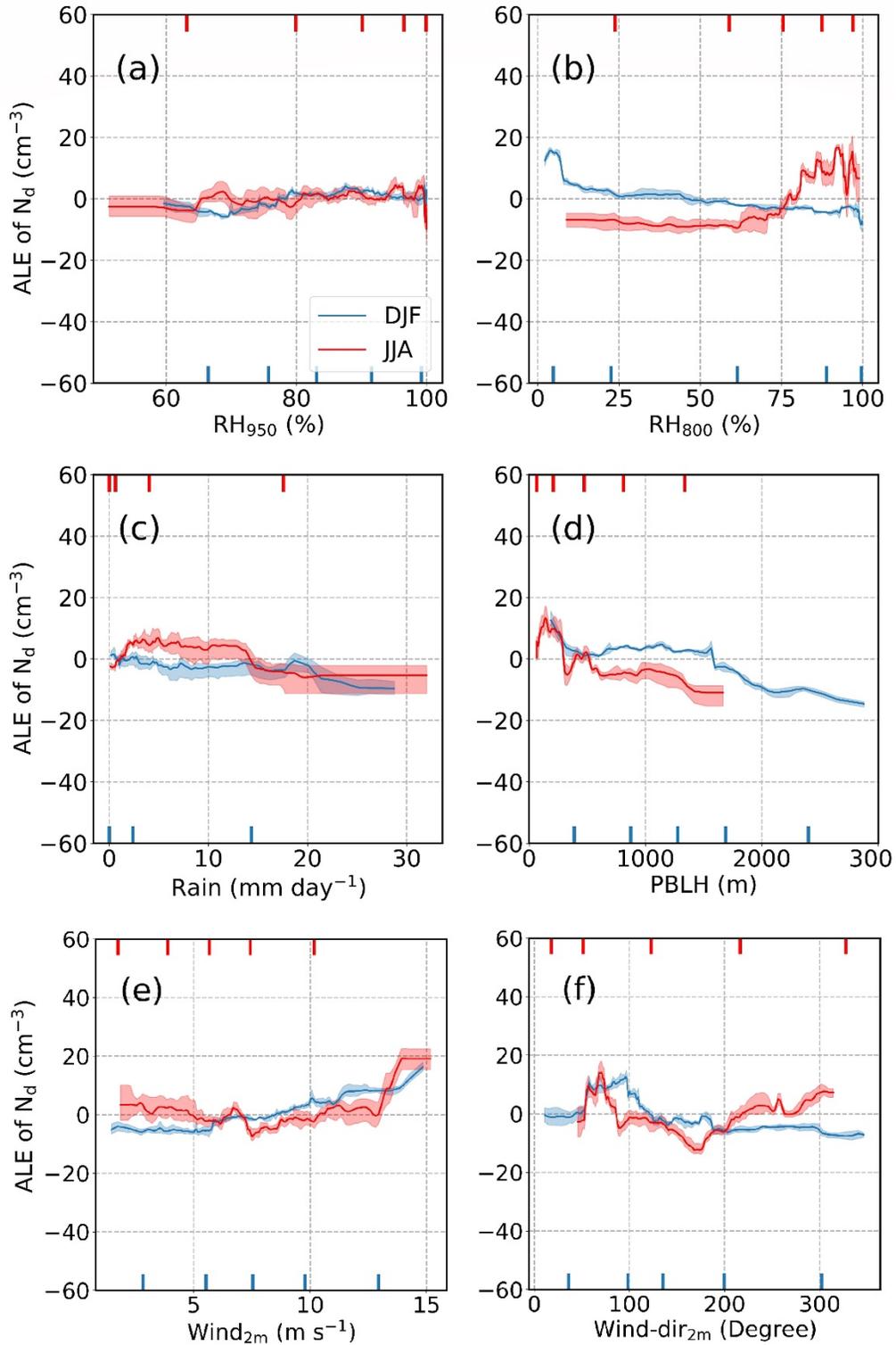
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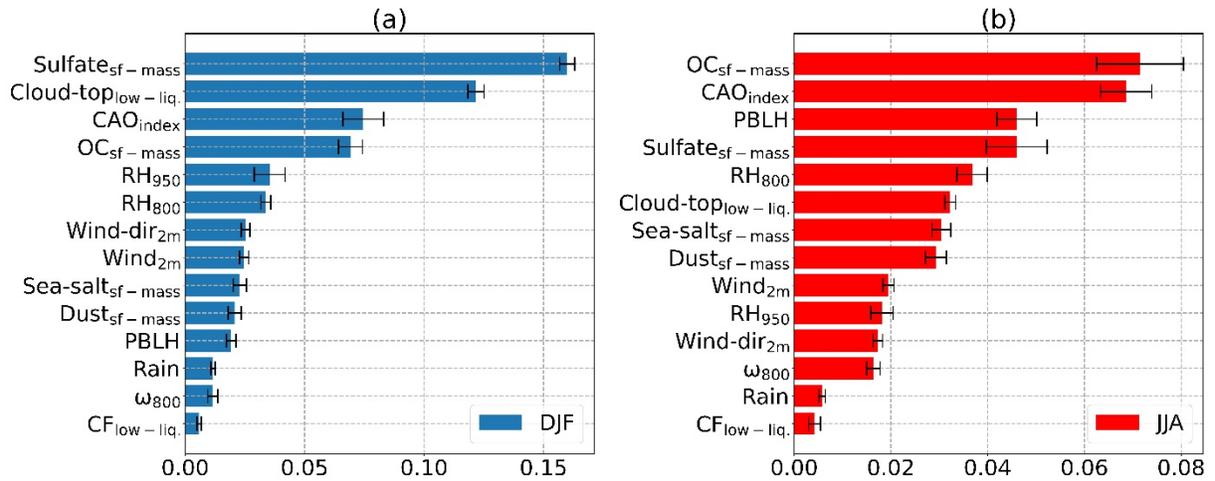
341 **Figure S24: Same as Figure S23 but for the following input parameters: (a) low-level liquid**
 342 **cloud fraction ($CF_{\text{low-liq.}}$), (b) cloud-top effective height of low-level liquid cloud ($\text{cloud-top}_{\text{low-}}$**
 343 **liq.), (c) cold-air outbreak (CAO) index, and (d) vertical pressure velocity at 800 hPa (ω_{800}).**

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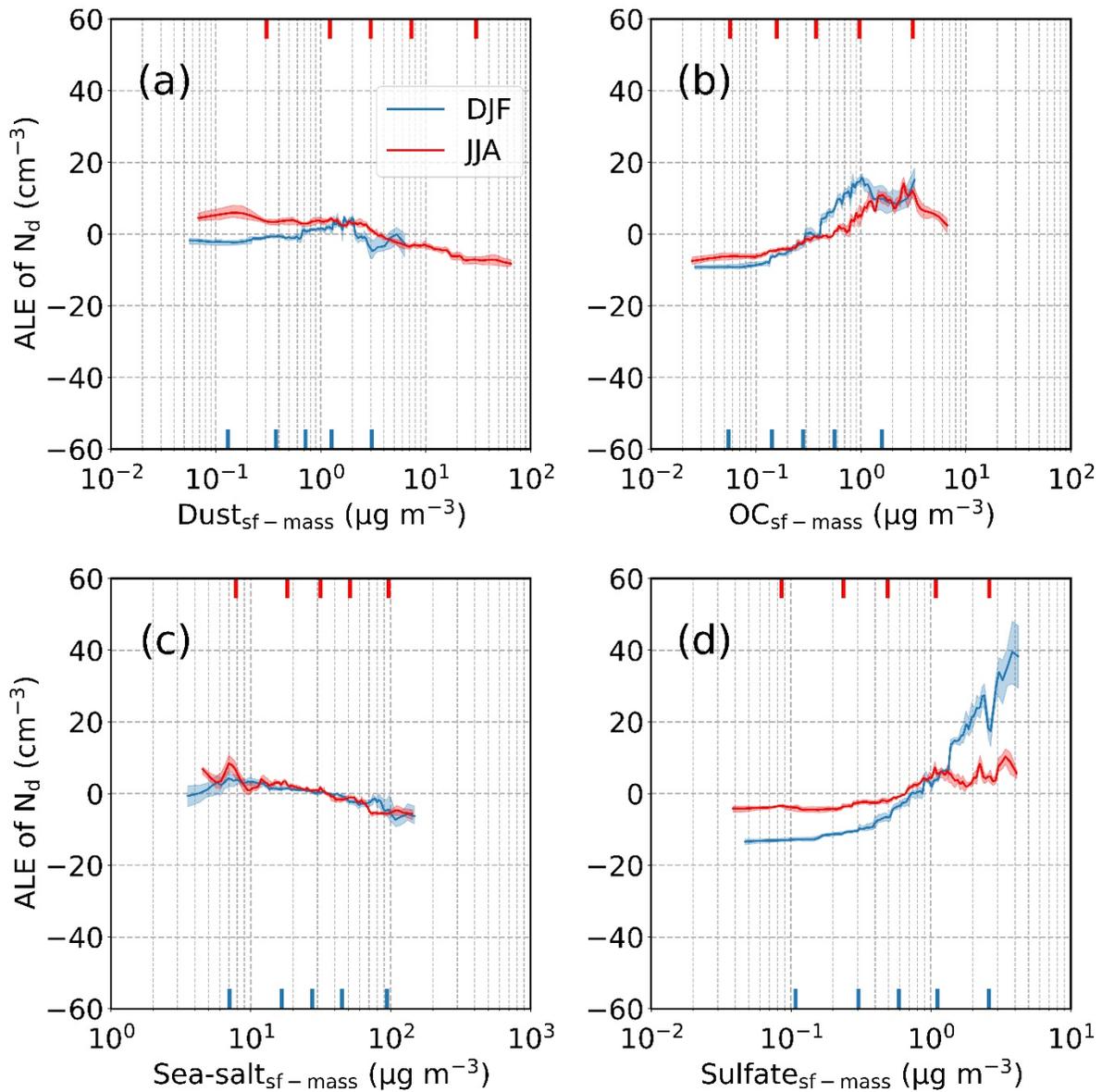
346 **Figure S25: Same as Figure S23 but for the following input parameters: (a) relative humidity**
 347 **at 950 hPa (RH_{950}), (b) relative humidity at 800 hPa (RH_{800}), (c) rain rate, (d) planetary**
 348 **boundary layer height (PBLH), (e) wind speed at 2 m (Wind_{2m}), and (f) wind direction at 2**
 349 **m (wind-dir_{2m}).**



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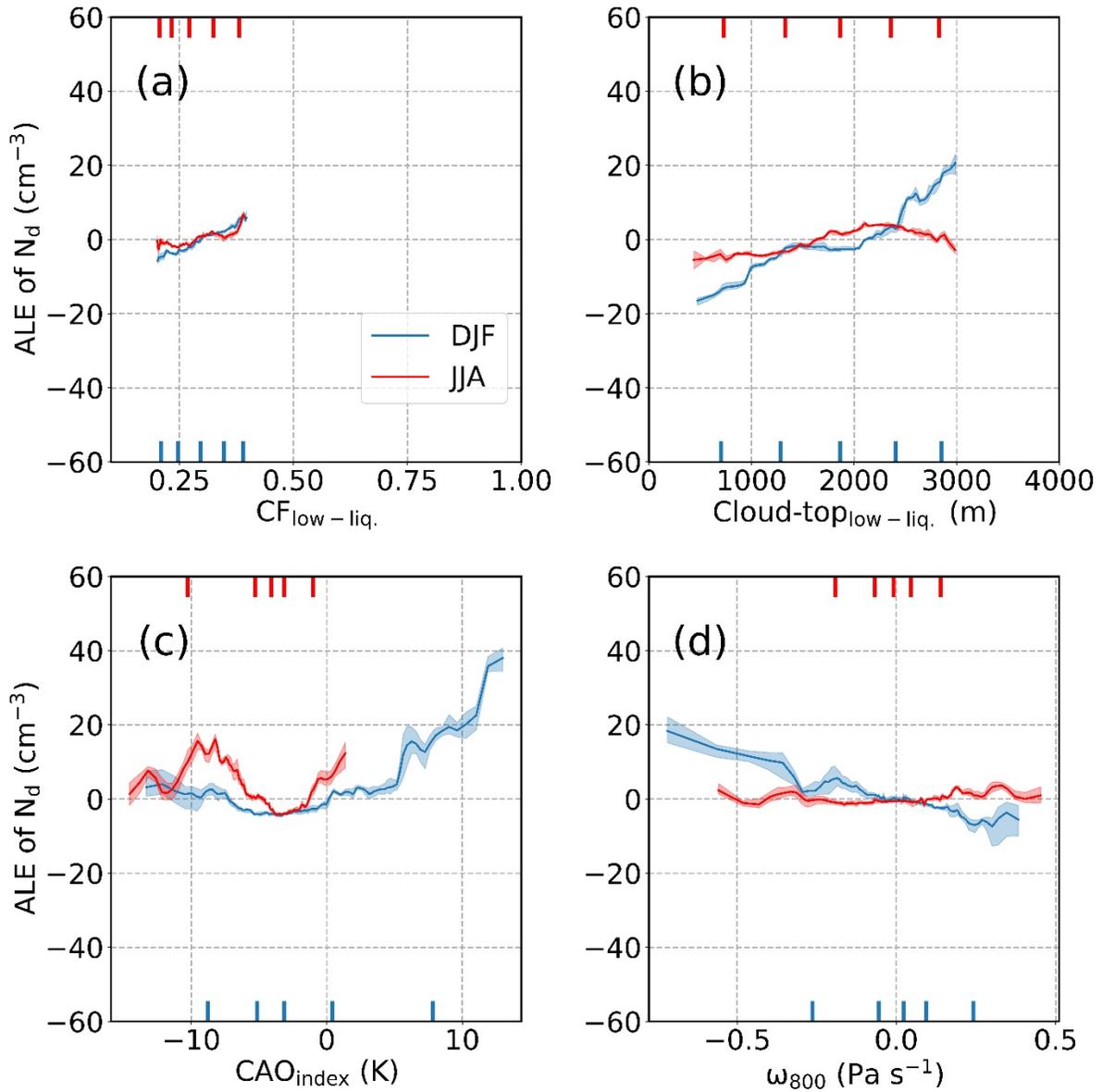
351 **Figure S26: Average permutation feature importance of input parameters for (a) DJF and**
 352 **(b) JJA based on GBRT models trained in each season on subsets of data including only**
 353 **samples with low-level liquid cloud fraction between 0.2 and 0.4 (i.e., $0.2 \leq CF_{low-liq.} \leq 0.4$).**
 354 **Feature importance values were calculated based on using the test set. Error bars exhibit the**
 355 **range of feature importance values stemming from the variability of the obtained models**
 356 **from the cross-validation resampling procedure.**

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 359 **Figure S27: Average local accumulated effect (ALE) profiles based on GBRT modeling for**
 360 **surface mass concentrations of the following parameters: (a) dust, (b) organic carbon, (c)**
 361 **sea-salt, and (d) sulfate. ALE profiles were based on GBRT modeling on subsets of data**
 362 **including only samples with low-level liquid cloud fraction between 0.2 and 0.4 (i.e., $0.2 \leq$**
 363 **$CF_{\text{low-liq.}} \leq 0.4$).** Blue and red profiles represent ALEs of DJF and JJA, respectively. Shaded
 364 **areas show the ALE ranges stemming from the variability of the obtained models from the**
 365 **cross-validation resampling procedure. Markers on the bottom and top x-axes denote the**
 366 **values of 5th, 25th, 50th, 75th, and 95th percentiles for each input variable.**

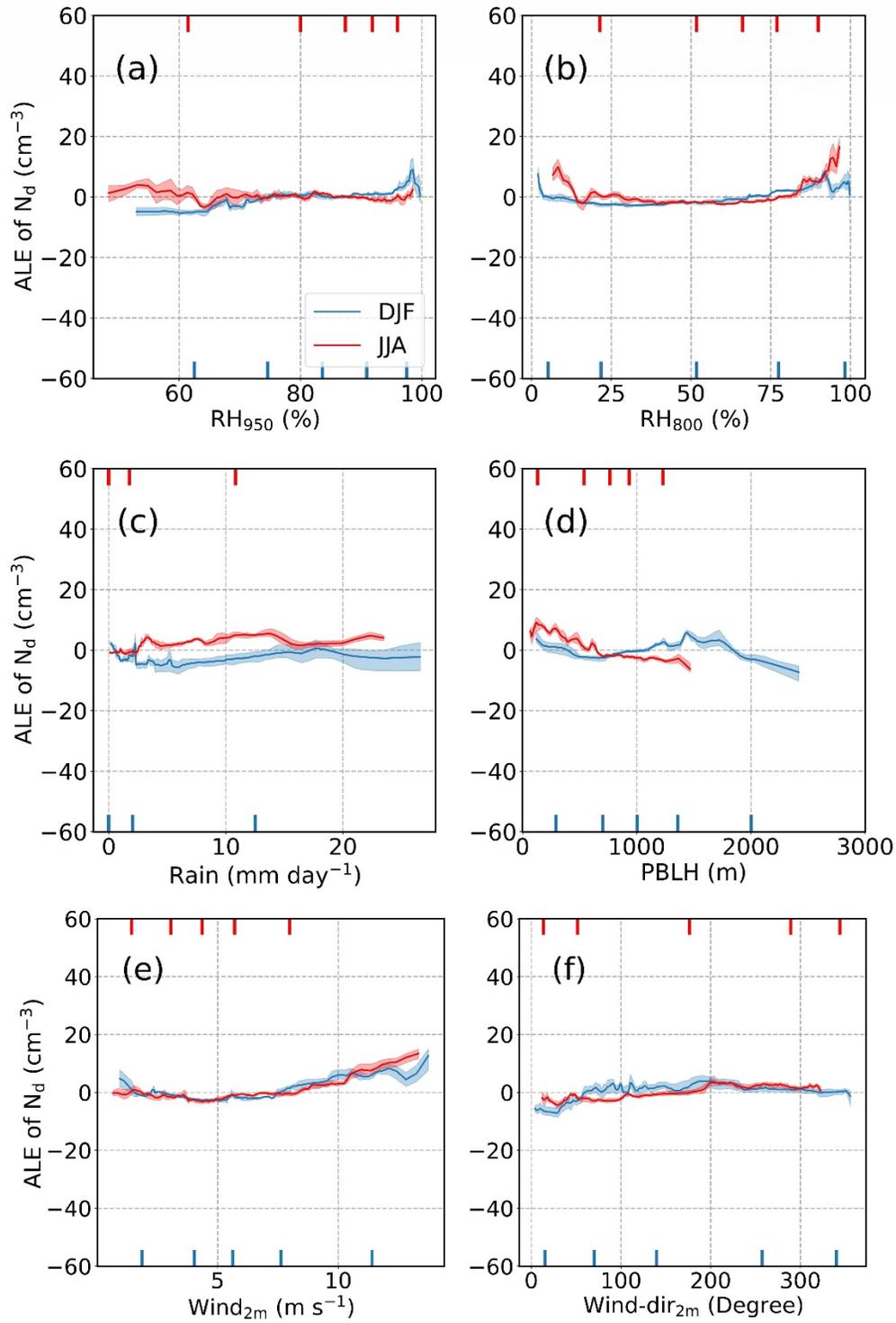
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368

369 **Figure S28: Same as Figure S27 but for the following input parameters: (a) low-level liquid**
 370 **cloud fraction ($CF_{\text{low-liq.}}$), (b) cloud-top effective height of low-level liquid cloud ($\text{cloud-top}_{\text{low-}}$**
 371 **liq.), (c) cold-air outbreak (CAO) index, and (d) vertical pressure velocity at 800 hPa (ω_{800}).**

372



373

374 **Figure S29: Same as Figure S27 but for the following input parameters: (a) relative humidity**
 375 **at 950 hPa (RH_{950}), (b) relative humidity at 800 hPa (RH_{800}), (c) rain rate, (d) planetary**
 376 **boundary layer height (PBLH), (e) wind speed at 2 m (Wind_{2m}), and (f) wind direction at 2**
 377 **m (wind-dir_{2m}).**

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- Could it be the case that the N_d retrievals don't work very well and give smaller N_d values when there are no boundary layer clouds (since they are designed to look at PBL liquid clouds)? This could then give higher N_d in the conditions with more PBL clouds (i.e., DJF). There is also likely to be more overlying higher altitude cloud in JJA, which would affect the retrieval of N_d . It would be good to look at the types of clouds and situations in which N_d is being retrieved in JJA – i.e., whether most of the retrievals come from mid-level clouds, or clouds with overlying ice cloud, etc.

386

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Response: We created a new Figure S1 to investigate the potential effects of N_d retrieval errors on the observed seasonal cycle of N_d . We also added some text at the end of section 3.2 to describe the results presented in Fig. S1.

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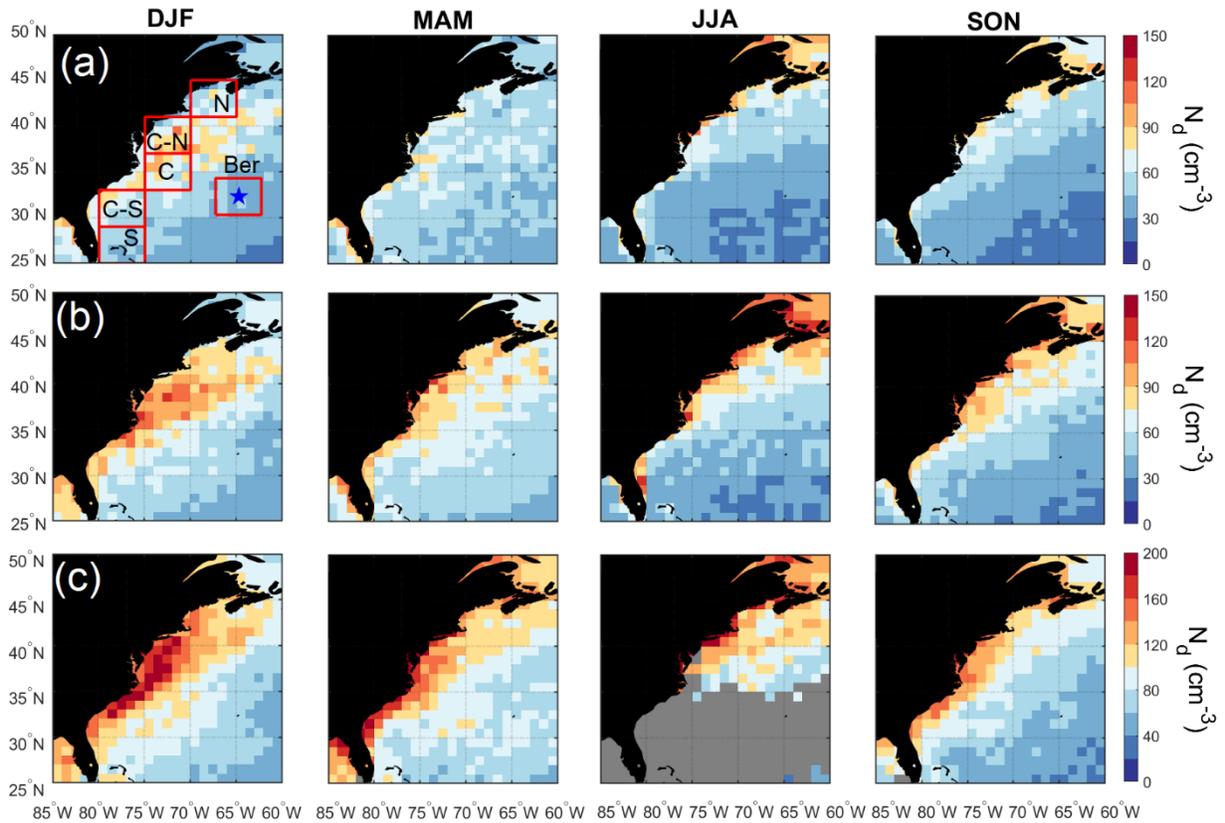
402

“One factor that could drive the seasonal variation in N_d is the unwanted effects of retrieval errors in the estimation of N_d at low cloud coverage conditions. Uncertainty associated with the estimation of N_d from MODIS observation increases as cloud fraction decreases (Grosvenor et al., 2018). This is mainly because of the overestimation of droplet effective radius (r_e) in the retrieval algorithm due to the interference of cloud-free pixels and also high spatial inhomogeneity in low cloud coverage conditions that violates horizontal homogeneity assumptions in the retrieval of r_e and τ from radiative transfer modeling (Zhang et al., 2012; Zhang et al., 2018). To test whether retrieval errors in N_d are the main driver of seasonal trends, Figure S1 shows the seasonal cycle of N_d at various low-level liquid cloud fractions. The results show that as cloud fraction increases the average N_d increases, regardless of season. Perhaps the more important result is that the seasonal trend in spatial maps of N_d remains similar regardless of cloud fraction. This finding is important as confirms that the seasonal cycle in N_d cannot be solely explained by the uncertainties associated with the retrieval of N_d at low cloud fraction.”

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407 **Figure S1: Seasonal maps of cloud drop number concentration for different ranges of low-**
 408 **level liquid cloud fraction ($CF_{\text{low-liq.}}$) as follows: (a) $0.1 \leq CF_{\text{low-liq.}} < 0.3$, (b) $0.3 \leq CF_{\text{low-liq.}} <$
 409 **0.6 , and (c) $CF_{\text{low-liq.}} \geq 0.7$. Gray pixels represent regions without sufficient sample points**
 410 **(less than 10 points) for calculating averages.****

411

412 Moreover, the reviewer raised a good point about the potential unwanted effects of high-
 413 altitude clouds in estimation of N_d . However, this should not be an issue for our analysis
 414 as we filter out clouds with cloud top pressure less than 700 hPa; thus, high altitude
 415 clouds were automatically removed from our analysis. We added the following
 416 information in section 2.1 to clarify this point:

417 “The CERES-MODIS SSF Level 3 product includes $1^\circ \times 1^\circ$ averaged data according to
 418 the cloud top pressure of individual pixels: low (heights below 700 hPa), mid-low (heights
 419 within 700–500 hPa), mid-high (heights within 500–300 hPa), and high (heights above 300
 420 hPa) level clouds. For this study, we only use low-cloud averages.”

- 421 • Can the composite analysis be done with the Cape Cod CCN data? Or can the CCN be
 422 used in the ML model? E.g., is it a better predictor than surface sulphate mass?

423

424 **Response: Unfortunately, it is not feasible to run the models with Cape Cod CCN data for**
 425 **two reasons. First, the data were only available for almost one year, which was not**

426 sufficient for our analysis. Second, and maybe even more important, CCN data was a
427 point measurement that provided data at most for only one grid cell over the WNAO
428 region ($25^{\circ} - 50^{\circ}\text{N}$ and $60^{\circ} - 85^{\circ}\text{W}$), which is not desirable for the analyses presented in
429 this work.

- 430 • Can you look at boundary layer decoupling index? Perhaps the high cloud fraction
431 regime in winter is more coupled than the summer regime allowing more efficient
432 transport of the surface aerosol to the clouds?
433

434 Response: We agree with the reviewer suggestion that the marine boundary layer in DJF
435 is more likely to be coupled than the conditions in JJA. This is supported by the results
436 presented in Painemal et al. (2021); they showed that surface heat fluxes including both
437 latent and sensible heat fluxes were substantially greater in DJF than JJA. In fact, JJA
438 exhibited the lowest surface heat fluxes among all the seasons, making it more likely to
439 have less turbulence and more prevalence of decoupled marine boundary layer
440 conditions. We addressed this comment by adding the following text at the end of section
441 3.5:

442 “The results of this section suggest though that aerosol indirect effects could be strongest
443 in DJF, meaning that N_d values increase more for the same increase in aerosol. Factors that
444 can contribute to higher ACI values in winter than summer include seasonal differences in
445 the following: (i) dynamical processes and turbulent structures of the marine boundary
446 layer; (ii) aerosol size distributions and consequently varying particle number
447 concentrations for a fixed mass concentration; and (iii) hygroscopicity of particles
448 especially as a result of changes in the composition of the carbonaceous aerosol fraction.
449 Regarding dynamical processes and the effects of turbulence, Figure 2 in Painemal et al.
450 (2021) shows that heat fluxes (i.e., latent and sensible fluxes) are strongest (lowest) in the
451 winter (summer) over the WNAO. The greater heat fluxes in DJF can contribute to more
452 turbulent and coupled marine boundary layer conditions in winter than summer,
453 presumably resulting in more efficient transport and activation of aerosol in the marine
454 boundary layer leading to higher ACI values. Forthcoming work will probe this issue in
455 greater detail.”

456

- 457 • It seems that the offshore flow prevalent in DJF (cf. southeasterlies from the open ocean
458 in JJA) might also play some role in determining N_d since it may transport more CCN
459 from the continent? Are there more measurements that can help to elucidate whether this
460 may be the case (e.g., aerosol size distribution data or additional CCN data from the Cape
461 Cod data)? It might be good to include the wind direction as a predictor in the ML?
462

463 Response: We already tried to answer this question by looking at Cape Cod data and
464 CALIPSO results presented in section 3.4. Another place where we looked at the effect of
465 size was the section 3.3 where we incorporated AI values rather than AOD in ACI
466 estimations.

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Based on this comment we also included wind direction at 2 m (Wind-dir_{2m}) as an input parameter of the GBRT models. Interestingly, we could not find a clear relationship between N_a and wind direction. We refer the reviewer to the previous comments to see the results of adding wind direction as an input parameter. This idea though is important and something that warrants additional investigation with more detailed data.

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- Other possible reasons for the DJF aerosols (or aerosols when there is lots of PBL cloud) being more efficient at making droplets could be discussed. E.g., could the aerosol be more hygroscopic in these conditions, could there be higher number concentrations (since here you mainly consider mass concentrations). These difference may be related to the different aerosol sources due to the different wind direction (see previous point).

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Response: We addressed this comment by adding the following text at the end of section 3.5: “The results of this section suggest though that aerosol indirect effects could be strongest in DJF, meaning that N_a values increase more for the same increase in aerosol. Factors that can contribute to higher ACI values in winter than summer include seasonal differences in the following: (i) dynamical processes and turbulent structures of the marine boundary layer; (ii) aerosol size distributions and consequently varying particle number concentrations for a fixed mass concentration; and (iii) hygroscopicity of particles especially as a result of changes in the composition of the carbonaceous aerosol fraction. Regarding dynamical processes and the effects of turbulence, Figure 2 in Painemal et al. (2021) showed that heat fluxes (i.e., latent and sensible fluxes) are strongest (lowest) in the winter (summer) over the WNAO. The greater heat fluxes in DJF can contribute to more turbulent and coupled marine boundary layer conditions in winter than summer, presumably resulting in more efficient transport and activation of aerosol in the marine boundary layer leading to higher ACI values. Forthcoming work will probe this issue in greater detail.”

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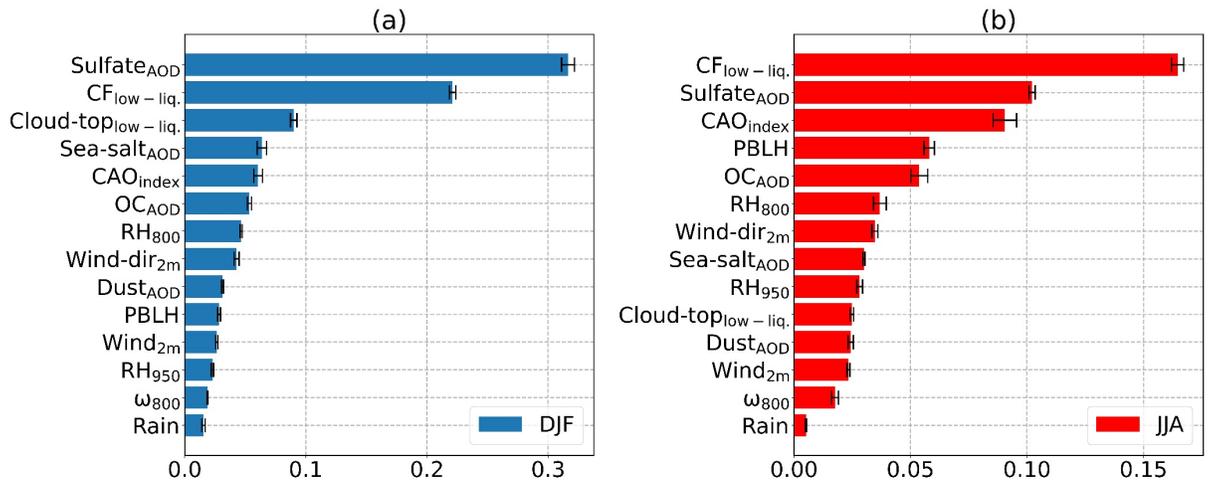
- Why did you not include AOD, speciated AOD, speciated boundary layer AOD, etc., in the ML model so that their impact relative to sulphate surface mass, etc. can be quantified?

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Response: We originally included these parameters in our early analyses but for simplicity (i.e., to cut down on having too many figures) we decided to only present the results for surface mass concentration as we thought the latter parameters were more relevant because they should be more closely linked to the aerosol level near the cloud base. We ran the model again though using speciated AOD values rather than surface mass concentrations. The general results are quite similar to the versions currently in the draft with some changes in the importance of parameters. Moreover, ALE plots are very similar to the versions in the main draft. As such, we decided to not include these results

508

in the main draft:

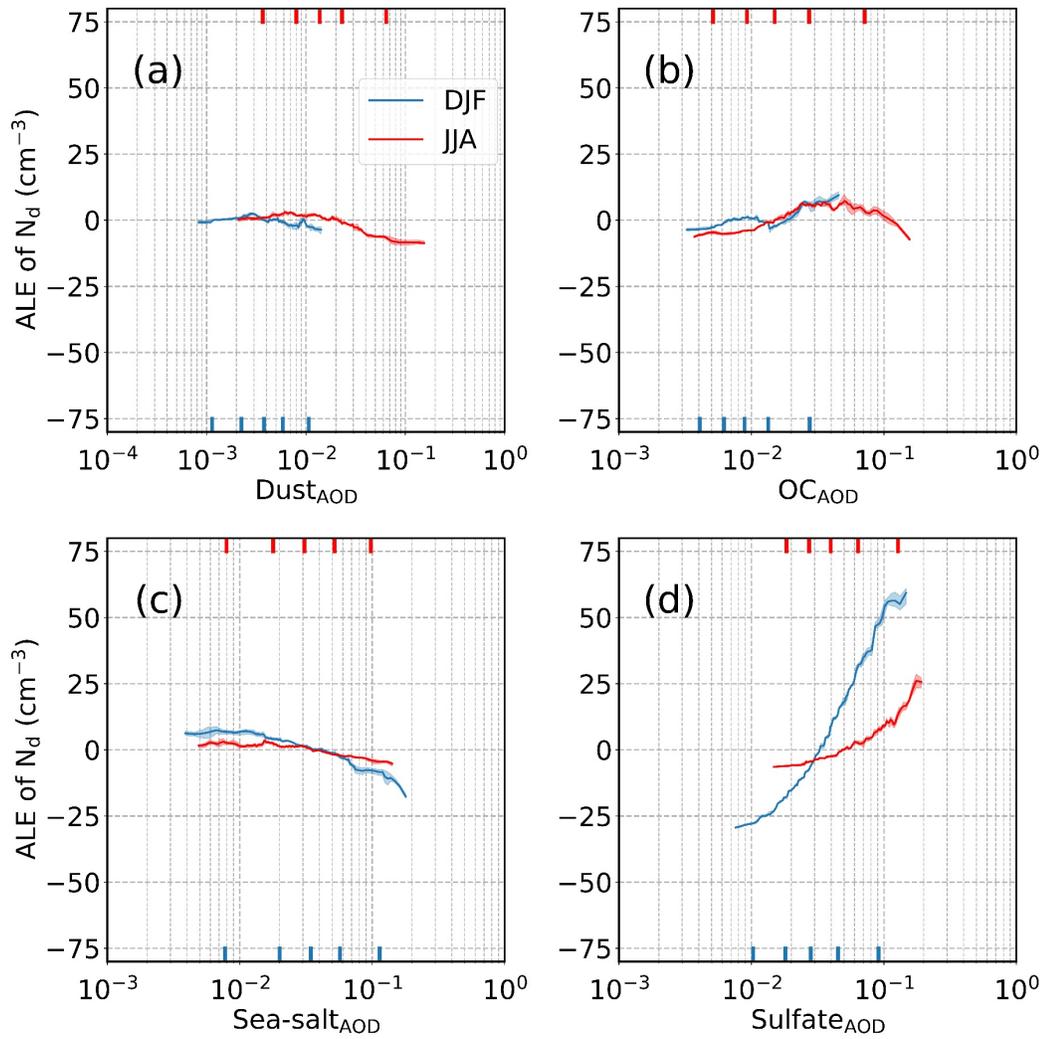


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Same as Figure 13 of the main draft but instead of surface mass concentrations, speciated AODs were included as input parameters to GBRT models.

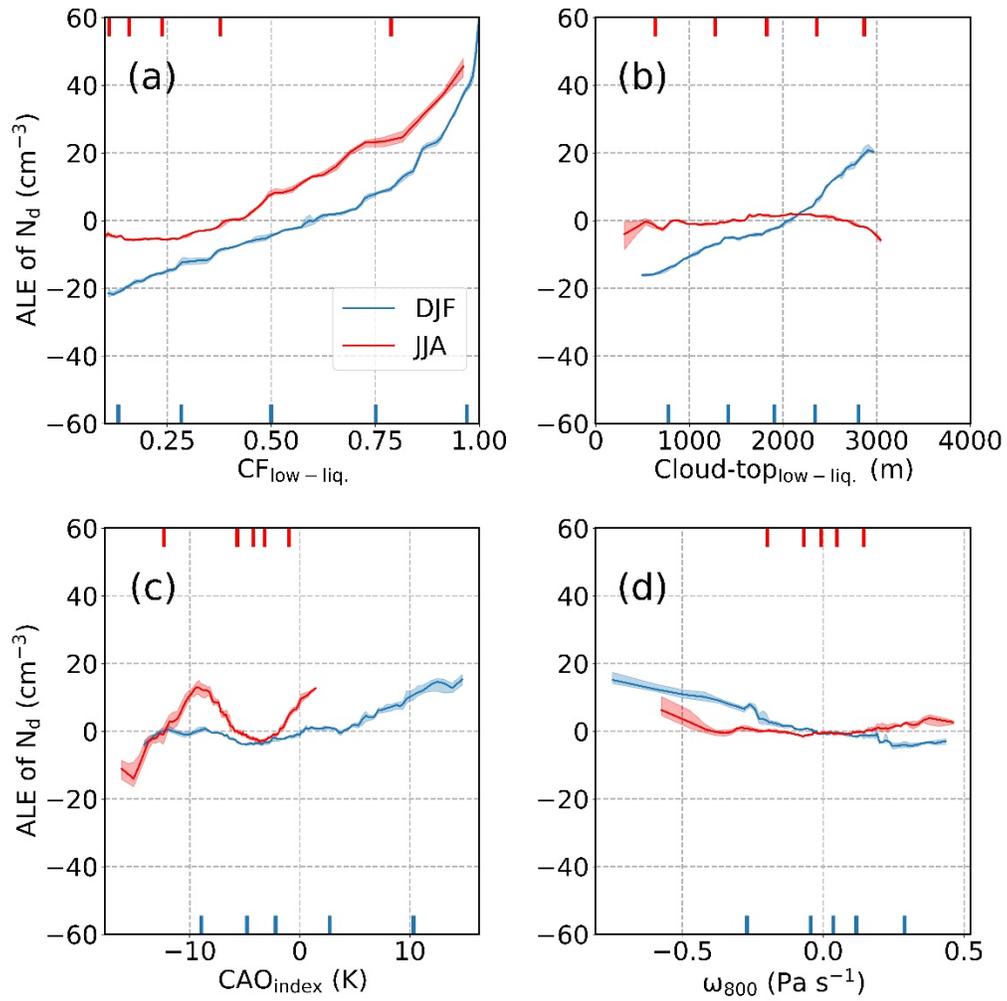


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Same as Figure 14 of the main draft but instead of surface mass concentrations, speciated AODs were included as input parameters to GBRT models.



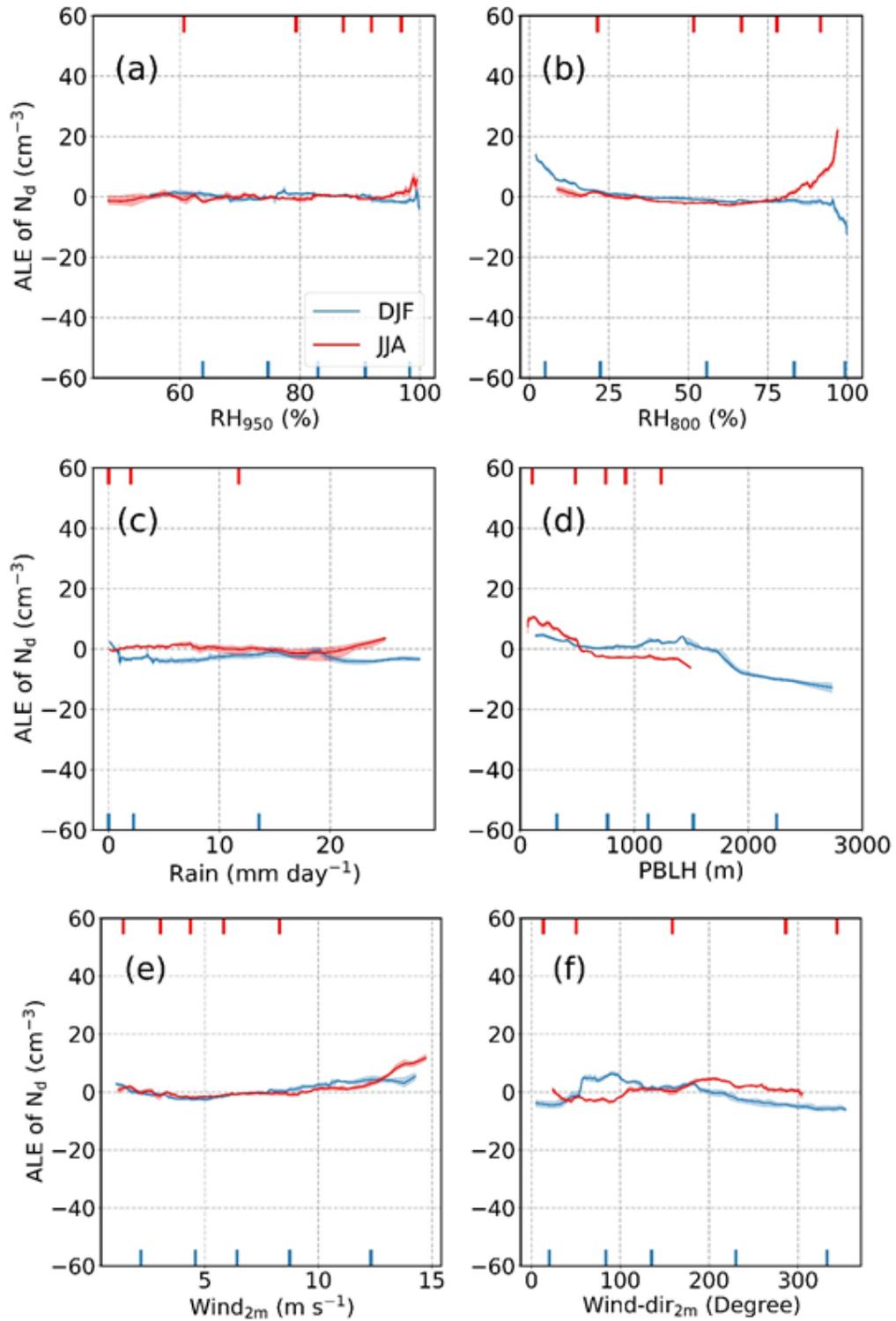
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Same as Figure 15 of the main draft but instead of surface mass concentrations, speciated AODs were included as input parameters to GBRT models.

518



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Same as Figure S21 of the main draft but instead of surface mass concentrations, speciated AODs were included as input parameters to GBRT models.

522

523

524 Also some more line specific comments :-

525

526 Section 3.1 – I'm not sure how well this section works where it is, or how useful the aircraft
527 analysis is for the main conclusions of the paper. Maybe it would be better placed at the end of
528 all the other results in order to help highlight some of the issues raised in the rest of the paper?

529

530 Response: We thought about this at great length when designing the very first draft we
531 originally submitted. We felt it was an exciting and compelling opener into the wide
532 gradient in N_d over the study region using a high quality (and “hard to get”) airborne field
533 dataset. We wanted to raise points from the airborne case study to motivate the general
534 topic we investigate throughout the rest of the draft. We still feel the section is suitable
535 where it currently is placed.

536

537 L318 – Is it possible to calculate an approximate activation diameter for 0.43%
538 supersaturation given the other aircraft measurements? This would make it easier to compare
539 to the $D_p > 10$ and $D_p > 3 \mu m$ data.

540 Response: This is a good idea but there are some challenges and overall we do not feel it is
541 needed for the overall story of the paper. More specifically, we would need to stitch together size
542 distribution data from various instrument like SMPS and LAS that measure different types of
543 diameters (e.g., aerodynamic diameter and electrical mobility diameter). The time resolution of
544 the SMPS is much longer than that of the LAS and, importantly, SMPS data were not available
545 in such a way to represent the whole flight segment under consideration.

546 L381 – “Consequently, humidity effects on remotely sensed aerosol parameters cannot alone
547 explain the dissimilar seasonal cycle of N_d and AOD, but can plausibly contribute to some
548 extent.”

549 - You haven't proved this quantitatively. Can you do a calculation of how much impact the RH
550 difference would have?

551

552 Response: Without knowing the exact composition of aerosol it is difficult to quantify the effects
553 of RH on extinction profiles. Therefore, we change the words to address the point reviewer
554 raised here:

555 “Consequently, humidity effects on remotely sensed aerosol parameters are less likely to be sole
556 explanation of the dissimilar seasonal cycle of N_d and AOD, but can plausibly contribute to some
557 extent.”

558

559

560 L411 – “but do not contribute significantly to number concentration as demonstrated clearly by
561 airborne observations (Figure 1).”

562 - It's not clear which part of Fig. 1 demonstrates this? Can you explicitly point this out?

563

564 Response: We addressed this comment by adding clarifying words:

565 “but do not contribute significantly to number concentration as demonstrated clearly by
566 airborne observations (i.e., FCDP_{>3μm} timeseries shown in Figure 1d).”

567 L413 – “This is supported in part by how DJF is marked by the highest fractional AOD
568 contribution from the PBL (59 – 72%) where sea salt is concentrated. In contrast, JJA has the
569 lowest fractional AOD contribution from the PBL (11.3 – 52.6%).”

570 - But this could also indicate higher CCN concentrations in the PBL in DJF perhaps due to the
571 aerosol being more trapped there than in summer?

572 Response: Here we only made a speculation, which we think is more likely given other
573 observations. To address the comment, we mention the other possibility based on the
574 reviewer suggestion:

575
576 “It is also possible that the higher fractional AOD contribution from the PBL in winter
577 partly owes to aerosol particles being more strongly confined to the PBL as compared to
578 the summer.”

579 Figure 3 – It looks like here the PBL AOD would be higher in DJF than for JJA for many of
580 the regions. It would be good to quote the PBL AOD values in Table 3. You need to describe
581 these results in more detail - the enhanced PBL values in DJF are an important point to
582 describe even though it seems that it doesn't explain the Nd seasonality.

583 Response: We already mentioned the percentage of AOD that comes from PBL and FT in
584 Table 3. We do not believe it is necessary to report these additional numbers.

585 L447 – “We next compare MERRA-2 speciated aerosol concentrations at the surface (Figure S2)
586 to those of speciated AOD (Figure S1).”

587
588 - I think that this information is very important and should not be in the supplementary
589 since whether the PBL aerosol mass concentrations (or ideally CCN number
590 concentrations, but I think they are not available from MERRA?) are lower in DJF
591 compared to JJA is a key part of the analysis regarding why the DJF AOD is lower and
592 yet the Nd higher compared to JJA. Indeed, Fig. S2 suggests that surface sulfate mass
593 concentration is about the same in DJF and JJA despite the AOD difference (and sulfate
594 AOD is also higher in JJA than DJF according to MERRA). Of course we might expect
595 sulfate mass concentrations to be even higher in DJF than in JJA if it was to explain the
596 Nd difference, but it does suggest that part of the issue is that AOD is vertically
597 integrated and is not just for the PBL.

598
599 Response: We made the requested change.

600
601 - Figs. S1 and S2 also suggest that near the coast sulfate AOD dominates over sea-salt,
602 which argues against the higher observed (Fig. 3) PBL extinction values in DJF being
603 due to sea-salt (as argued e.g., L413 and L707).

604 Response: A closer inspection of the former Figure S2 though shows very different
605 number scales for sea salt versus sulfate that should not be ignored. As that figure

606 quantifies surface mass concentrations (relevant to PBL), we feel it still supports sea
607 being a dominant contributor to PBL aerosol mass and optical depth.

608 - Although sea-salt surface mass concentrations are higher than sulfate. Do you have
609 speciated profiles of extinction from MERRA? These could be used to quantify the effect
610 of sea-salt on the PBL AOD.

611 Response: We looked for such data but could not locate it as part of the standard
612 MERRA-2 data products.

613 - What seems a bit strange given that DJF and JJA have similar surface SO₄ in MERRA
614 is that the Cape Cod observations show lower CCN concentrations in DJF. It would also
615 be good to discuss this a bit more along with potential caveats. You mention that the 1%
616 supersaturation at which the CCN are measured is quite high and would be counting
617 fairly small aerosol particles – I think it would be good to show the lower
618 supersaturations that you say are available. Or at least check whether the DJF values of
619 these are also lower than in JJA (data permitting). Is there any other CCN data down the
620 east coast of US since it would be very useful here. Or do you have observed aerosol size
621 and composition measurements that might help determine whether the supersaturation
622 has a big effect and whether there really are fewer CCN in DJF?

623
624 Response: Regarding CCN data at lower supersaturation, unfortunately the temporal
625 coverage of CCN data at Cape Cod was not good enough to give a full seasonal profile at
626 lower supersaturation. About other datasets along the East Coast we do not have those at
627 our disposal to use. These are all good avenues of future research.

628 o Also, how representative is the 1-year of data likely to be? Could the
629 interannual variability be large enough to make that result uncertain?

630
631 Response: This is a hard question to robustly answer without any uncertainty without
632 having the data. It is our anticipation that there would not be significantly high year-to-
633 year variability to change the story. We did not feel this comment needed additional
634 revision.

635
636 o It could also be that MERRA is doing a poor job of representing the sulphate
637 mass concentration.

638
639 Response: Answering such a question is outside to scope of our study.

640 o Finally, it would be good to quantify how likely it is that we can have a similar
641 surface sulphate mass concentration, but different CCN (using observations or a
642 more sophisticated model perhaps).

643

644 Response: It is our opinion that this type of question is more geared towards future work
645 and outside the scope of this current paper’s objectives.

646 Typos / grammar etc.

647 L252 – “The ALE value of feature S” – it’s not clear here what you mean by “feature”. The
648 symbols for this equation have also not all been explained. What does subscript c refer to? What
649 are x_s , x_c , z_s ? “The value of $f_{s,ALE}(x_s)$ can be viewed as the difference between the model’s
650 response at x_s and the average prediction.” – this is not very clear – I think this explanation
651 needs to be clearer for interpreting the later figures.

652

653 Response: We clarified these variables as follows:

654 “ALE plots illustrate the influence of input variables on the response parameter in ML
655 models. The ALE value for a particular variable S at a specific value of x_s (i.e., $f_{s,ALE}(x_s)$)
656 can be calculated as follows:

$$657 \quad f_{s,ALE}(x_s) = \int_{z_{0,1}}^{x_s} \int_{x_c} f^s(z_s, x_c) P(x_c|z_s) dx_c dz_s - constant \quad (2)$$

658 where $f^s(z_s, x_c)$ is the gradient of model’s response with respect to variable S (i.e., local
659 effect) and $P(x_c|z_s)$ is the conditional distribution of x_c where C denotes the other input
660 variables rather than S and x_c is the associated point in the variable space of C. $z_{0,1}$ is chosen
661 arbitrarily below the smallest observation of feature S (Apley and Zhu, 2020).”

662

663 L288 – the flight leg abbreviation “BCT1” is stated here without definition. It would be good to
664 have introduced the different flight legs being described before this with reference to Fig. 1.

665

666 Response: We added a line referring readers to the caption of Fig.1 for the definition of
667 different legs: “Sea surface temperatures were 6 – 9°C near the coast during the descent
668 and Min. Alt. 1 leg (readers are referred to Fig. 1’s caption for the definition of different
669 legs)...”

670 L305 – “ranged” -> “ranged from”.

671

672 Response: Fixed as follows: “ N_d values from the FCDP ranged from a maximum value of
673 1298 cm^{-3} ...”

674 L307 – “which is a fairer comparison with the ACB1 leg” – fairer than what?

675

676 Response: Fixed as follows: “The minimum N_d value in the ACB3 leg was 85 cm^{-3}
677 (34.11° N, 72.80° W), which is a fairer comparison to the ACB1 leg as compared to the
678 BCT1 leg in terms of being closer to cloud base.”

679 L311 – “there was a significant offshore gradient in LAS submicrometer particle number
680 concentration and AMS non-refractory aerosol mass, ranging from 424 cm^{-3} and 5.60 $\mu\text{g m}^{-3}$
681 (from BCB1) to 21 cm^{-3} and 0.32 $\mu\text{g m}^{-3}$ (from BCB3), respectively; these values are based on

682 times of the maximum and minimum LAS concentrations during the BCB1 and BCB3 legs,
683 respectively.”

684 - This sentence was a little confusing and could be made clearer I think.

685

686 Response: We revised the text: “there was a significant offshore gradient in LAS submicrometer
687 particle number concentration and AMS non-refractory aerosol mass, ranging from as high as
688 424 cm^{-3} and $5.60 \mu\text{g m}^{-3}$ (during BCB1) to as low as 21 cm^{-3} and $0.27 \mu\text{g m}^{-3}$ (during BCB3).”

689

690 L320 – “There was a slighter gradient in particle concentrations with $D_p > 3 \mu\text{m}$ (most likely sea
691 salt) between the same two points of maximum and minimum LAS concentration in BCB1 and
692 BCB3 legs, respectively: 0.26 cm^{-3} to 0.11 cm^{-3} .”

693 - This could be written more clearly.

694

695 Response: Revised this general section for clarity: “For the duration of the flight portion shown
696 in Figure 1, supermicrometer concentrations varied over two orders of magnitude ($0.002 - 0.51$
697 cm^{-3}) and expectedly did not exhibit a pronounced offshore gradient as it is naturally emitted
698 from the ocean.”

699

700 Fig. 4 – “The notches in the box plots demonstrate whether medians are different with 95%
701 confidence.”

702 - Different to what? Or do you mean it shows the 95% confidence range of the median?

703

704 Response: They show whether medians are different from each other with 95% confidence. We
705 updated the caption as follows: “**The notches in the box plots demonstrate whether medians**
706 **are different from each other with 95% confidence. Boxes with notches that do not overlap**
707 **with each other have different medians with 95% confidence.**”

708

709 L486 – “Coefficients of determination (R^2) when computing seasonal ACI values”

710 - What do you mean by this? Is this the correlation coefficient between N_d and the
711 aerosol proxy?

712

713 Response: We clarified this as follows: “Coefficients of determination (R^2) for the linear
714 regression between $\ln(N_d)$ and $\ln(\alpha)$ when computing seasonal ACI values were generally low (\leq
715 0.30), with spatial maps of R^2 and data point numbers in Figure S2.”

716

717 L521 – “Subsequently, one standard deviation from both sides of the seasonal mean defined cut-
718 off points outside of which we assign values as being low and high in each season.”

719 - Could be written more clearly.

720

721 Response: Revised text: “We assign values as being low in each season if they are less
722 than one standard deviation below the seasonal value; conversely, high values are those
723 exceeding one standard deviation above the seasonal mean.”

724
725 L683 – “will struggle for analysing”.

726
727 Response: Revised text: “Satellite remote sensing studies of aerosol-cloud interactions
728 presumably will be more challenging in winter periods versus the summer with regard to
729 the spatial and temporal mismatch between cloud and aerosol retrievals.”

730 L702 – “There were significant changes” – what were the changes?

731
732 Response: We added the following lines in response to this comment: “These changes
733 included a sharp decrease in aerosol number concentration, a decrease in mass fraction of
734 sulfate in droplet residual particles, and an increase in mass fraction of organic and
735 chloride of droplet residual particles moving offshore.”

736
737 L705 – “and surface-based aerosol mass concentrations and CCN concentrations (1%
738 supersaturation) are generally highest in JJA and MAM and are at (or near) their lowest
739 values in DJF”

740 - Surface sulfate aerosol mass concentrations were actually similar in DJF and JJA.

741
742 Response: Revised text: “ N_d is generally highest (lowest) in DJF (JJA) over the WNAO
743 but aerosol parameters such as AOD, AI, surface-based aerosol mass concentrations for
744 most species, and CCN concentrations (1% supersaturation) are generally highest in JJA
745 and MAM and are at (or near) their lowest values in DJF.”

746
747 L725 – “by CAO type of conditions” – better as “by conditions associated with CAOs”.

748
749 Response: Fixed: “...which is assisted in large part by conditions associated with CAOs
750 such as high cloud fraction and high CAO index.”

751
752
753 References:

754 Painemal, D., Corral, A. F., Sorooshian, A., Brunke, M. A., Chellappan, S., Goroooh, V. A., Ham,
755 S., O’Neill, L., Smith Jr., W. L., Tselioudis, G., Wang, H., Zeng, X., and Zuidema, P.: An
756 Overview of Atmospheric Features Over the Western North Atlantic Ocean and North American
757 East Coast – Part 2: Circulation, Boundary Layer, and Clouds, J Geophys Res-Atmos,
758 10.1029/2020JD033423, 2021.

759