Response to Comments of Reviewer #1

(comments in *italics*)

Manuscript number: acp-2021-119

Title: Impacts of aerosol-photolysis interaction and aerosol-radiation feedback on surface-layer ozone in North China during a multi-pollutant air pollution episode

Yang et al. examined the impacts of aerosols on surface ozone through the two well-known pathways, i.e., aerosol-photolysis interaction and aerosol-radiation feedback. The novelty of this study is its focus on the polluted episodes with elevated both $PM_{2.5}$ and ozone levels over North China. They also quantified the chemical and physical processes that drive the aerosol-radiation interactions.

Overall, this is a timely study and it clearly demonstrates the impacts of aerosols on ozone pollution. The structure of this manuscript is easy to follow. Although some of the manuscript needs further clarification, the results are generally convincing. As such, I think it is publishable after the following issues are addressed.

Response:

Thanks to the reviewer for the valuable comments and suggestions which are very helpful for us to improve our manuscript. We have revised the manuscript carefully, as described in our point-to-point responses to the comments.

Specific Comments:

1. *In Abstract: ozone changes refer to MDA8 ozone or daytime ozone?*

Response:

The ozone changes in abstract is daytime ozone. According to the reviewer's comments, we have added this information in the revised manuscript. (**Page 2, Line 33**)

2. *Line 177: a correlation coefficient of 0.66 reads like not high!*

Response:

According to the reviewer's suggestion, we have corrected it in the revised manuscript as follows: "The model can also reasonably capture the temporal variations of observed $PM_{2.5}$ and O_3 with correlation coefficients (R) of 0.66 for $PM_{2.5}$ and 0.86 for O_3." (**Page 7, Line 180-183**)

3. *Lines 179-181: the oxidation of SO_2 by NO_2 in aqueous aerosols is important for summertime?*

Response:

Thanks for your suggestion, we have changed the explanation in the revised manuscript as follows: "The failure to reproduce $PM_{2.5}$ peak values may be attributed to incomplete treatments of chemical reactions in WRF-Chem, e.g., missing the heterogeneous chemistry in the model (Cheng et al., 2016) and the lack of secondary organic aerosols (SOA) formation pathways in the aerosol module (Chen et al., 2019)." (**Page 7-8, Line 183-188**)

4. *Lines 248-251: this statement looks reasonable here, but in the later text the process analysis shows that chemistry will be enhanced by ARF. Instead, ARF decreases ozone through physical processes.*

Response:

Thanks for the reviewer's suggestion. We have deleted this sentence in the revised manuscript.

5. *Line 260: "is" should be "are". Please do proof-reading throughout the text.*

Response:

This sentence has been deleted in the revised manuscript. According to the reviewer's comments, proof-reading has been conducted through the whole revised manuscript.

6. *Line 310: It is Okay to use model levels (e.g., 12 levels), but it will be better to add model height in meters as well.*

Response:

Thanks for reviewer's suggestion. We have added the model height in meters in the revised manuscript. (**Page 12, Line 314, Line 317-319, Line 328**)

7. *Lines 326-327: why do you need this statement?*

Response:

Analyzing Fig. 8c we can conclude that ARF promotes the O_3 chemical production with a positive mean value of 0.66 ppb h^{-1}. The enhanced O_3 precursors due to ARF can promote the chemical production of O_3. According to the reviewer's comment, we have deleted this statement in the revised manuscript.

8. *Lines 327-328: Please provide evidence to support this conclusion.*

Response:

The typical VOCs/NO_x ratio is calculated to classify sensitivity regimes and to indicate the possible O_3 responses to changes in VOCs and/or NO_x concentrations. O_3 production is VOC-limited if the ratio is less than 4, and it is NO_x-limited if the ratio is larger than 15 (Edson et al., 2017; Li et al., 2017). The ratio of VOCs/NO_x ranging around 4-15 indicates a transitional regime, where ozone is nearly equally sensitive to each species (Sillman, 1999). As shown in Fig. R1(a-c), O_3 are mainly formed under VOC-limited and transition regimes in CAPAs, which means that the increased concentrations of VOCs and NO_x are favorable for ozone chemical production. As shown in Fig. R1(e) and (h), both the surface concentrations of VOCs and NO_x are increased when the impacts of ARF are considered. Thus, the contribution of CHEM in NOAPI is larger than that in NOALL. Similar results can also be found in Gao et al. (2018).

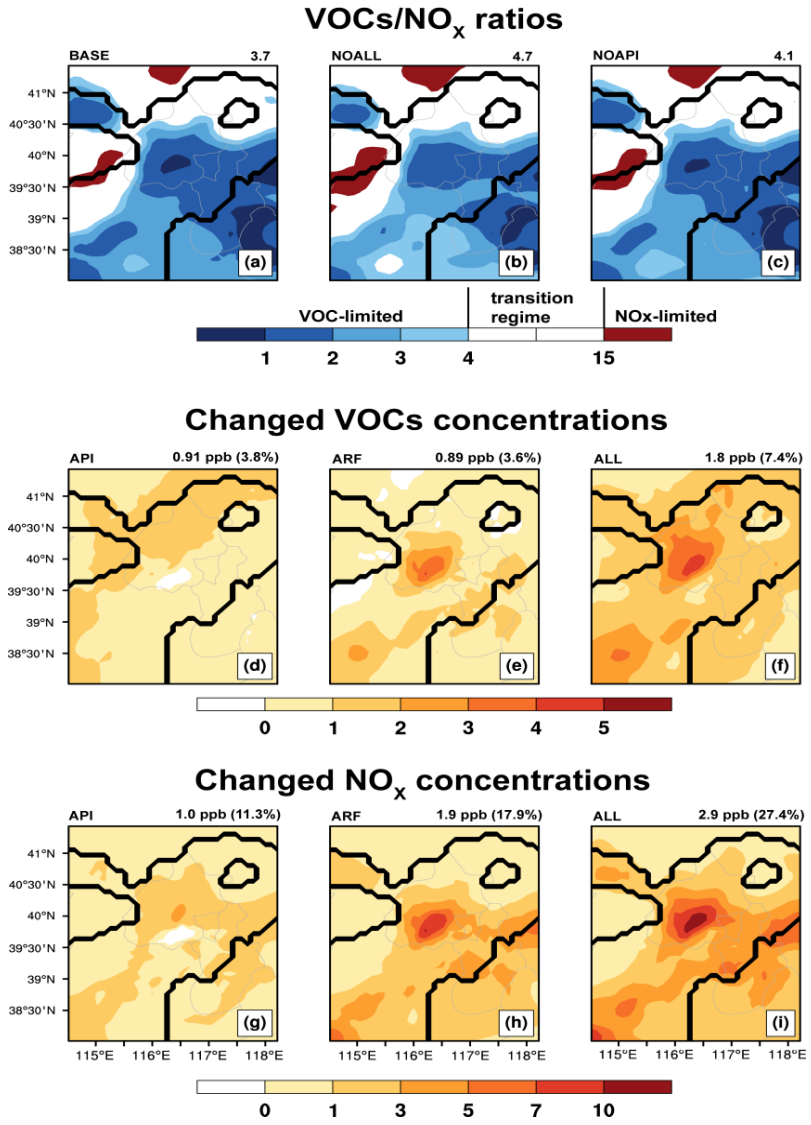

Figure R1. The ratios of VOCs/NO_x calculated from (a) BASE, (b) NOALL, and (c) NOAPI. The changed surface-layer concentrations of VOCs and NO_x (NO_2+NO, ppb) caused by (d, g) API, (e, h) ARF, and (f, i) ALL during the daytime (08:00-17:00 LST) from 28 July to 3 August 2014. The calculated values averaged over CAPAs are also shown at the top of each panel.

9. *Discussion. I think the authors should do some comparisons between your results with previous studies. This is important for readers to better understand your case study results. Moreover, how about the applicability of the calculated ROP of -0.14 ppb ($\mu g\ m^{-3}$)$^{-1}$?*

Response:

According to the comments of Reviewer#2, another two complex air pollution episodes (8-13 July 2015 and 5-11 June 2016) in this region are selected to conduct simulations for generating general conclusions (**Page 13-14, Line 343-365**). Meanwhile, a discussion about the impacts of secondary organic aerosols (SOA) is also added in the section 6 (**Page 15, Line 402-412**).

Thanks to the reviewer's comments. As the relationship between O_3 and $PM_{2.5}$ is non-linear, and the simple index of ROP can not fully represent the impacts of aerosols on surface O_3, so we

delete the ROP in the revised manuscript.

10. Fig.2: It will be better to add error bars for observed $PM_{2.5}$ and ozone.

Response:

According to the reviewer's suggestion, error bars have been added in Fig. 2 in the revised manuscript. (**Page 27**)

11. Fig.3: what are the cities these plots for?

Response:

The averaged T_2, RH_2, and WS_{10} are collected from ten meteorological observation stations, and the detail information about the sites is listed in Table S1. The photolysis rates of NO_2 ($J[NO_2]$) are observed in Peking University. More details are explained in section 2.3. (**Page 6**)

12. Fig.7: what are the layers your process analysis applied for? I don't see this key information here, as well as in the text.

Response:

The surface-layer, namely, first-layer O_3 concentrations are analyzed in Fig. 7. Thanks for reviewer's suggestion, we have added this information in the revised manuscript. (**Page 10, Line 266**)

Reference:

Chen, L., Zhu, J., Liao, H., Gao, Y., Qiu, Y., Zhang, M., Liu, Z., Li, N., and Wang, Y.: Assessing the formation and evolution mechanisms of severe haze pollution in the Beijing–Tianjin–Hebei region using process analysis, Atmos. Chem. Phys., 19, 10845–10864, https://doi.org/10.5194/acp-19-10845-2019, 2019.

Cheng, Y., Zheng, G., Chao, W., Mu, Q., Bo, Z., Wang, Z., Meng, G., Qiang, Z., He, K., and Carmichael, G.: Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events in China, Science Advances, 2, https://doi.org/10.1126/sciadv.1601530, 2016.

Edson, C. T., Ivan, H.-P. and Alberto, M.: Use of combined observational- and model-derived photochemical indicators to assess the O_3-NO_x-VOC System sensitivity in urban areas, Atmosphere., 8, 22. https://doi.org/10.3390/ atmos8020022, 2017.

Li, K., Chen, L., Ying, F., White, S. J., Jang, C., Wu, X., Gao, X., Hong, S., Shen, J., Azzi, M. and Cen, K: Meteorological and chemical impacts on ozone formation: a case study in Hangzhou, China, Atmos. Res., 196, https://doi.org/10.1016/ j.atmosres.2017.06.003, 2017.

Sillman, S.: The relation between ozone, NO_x and hydrocarbons in urban and polluted rural environments, Atmos. Environ., 33, 1821-1845, https://doi.org/ 10.1016/S1352-2310(98)00345-8, 1999.

Thank you very much for your comments and suggestions.

Response to Comments of Reviewer #2

(comments in *italics*)

Manuscript number: acp-2021-119

Title: Impacts of aerosol-photolysis interaction and aerosol-radiation feedback on surface-layer ozone in North China during a multi-pollutant air pollution episode

In this study, Yang et al. investigate the impact of aerosol-radiation interactions on O_3 formation during a multi-pollutant air pollution episode in Northern China. Additionally, the study uses process analysis to analyze how the aerosol-radiation interactions affect O_3 through various physical and chemical mechanisms. This is an interesting research topic with valid research methods and an overall well written and well-structured manuscript. However, the period of analysis is far too short (i.e., 7 days) to robustly quantify the impact of aerosol-radiation impacts in this region or to describe any variability. Additionally, the time period analyzed appears somewhat arbitrary and is nearly a decade removed from current conditions. For these reasons, the manuscript is not currently at the scientific level of the Atmospheric Chemistry and Physics Journal. However, this manuscript would be suitable for publication in ACP if either it is restructured to focus on how the methods used are unique and different from past work or if the authors investigate longer periods to generate more robust analysis and conclusions. Please find my major and minor comments below.

Response:

Thanks to the reviewer for the valuable comments and suggestions which are very helpful for us to improve our manuscript. We have revised the manuscript carefully, as described in our point-to-point responses to the comments.

The major innovation of this study is that it is the first time to quantify the respective/combined contributions of aerosol-photolysis interaction (API) and aerosol-radiation feedback (ARF) on O_3 concentrations during multi-pollutant air pollution episodes characterized by high O_3 and $PM_{2.5}$ levels. According to the reviewer's comments, another two complex air pollution episodes are also analyzed for generating general conclusions, and we find that API is the dominant factor for O_3 reduction related to aerosol-radiation interactions during all the simulated episodes (Episode 1: 28 July-3 August 2014; Episode 2: 8-13 July 2015; Episode 3: 5-11 June 2016).

Major Comments:

1. *The novelty of this study is that it is the first time that API and ARF are investigated for synchronous occurrences of high $PM_{2.5}$ and O_3 concentrations. This is a rather broad research question to be focused on only one region and one very minor time period. Why do the authors not conduct simulations for either several of these small pollution episodes in this region or for similar episodes in other locations in China?*

Response:

The high-resolution WRF-Chem model has been widely applied to investigate the evolution mechanisms of air pollutants during short time periods (Gao et al., 2016; Qiu et al., 2017; Gao et

al., 2018; Wang et al., 2020). Gao et al. (2016) summarized the general conclusion that haze events were mainly caused by high emissions of air pollutants and unfavorable weather conditions in North China Plain (NCP) by analyzing a simulated pollution episode from WRF-Chem during 14-24 January 2010. According to the results from WRF-Chem, Qiu et al. (2017) reported that the direct radiative effects of scattering aerosols were greater than that of absorbing aerosols in NCP during 21-27 February 2014. Gao et al. (2018) found that the interactions between black carbon and planetary boundary layer (PBL) could influence the surface O_3 concentration in Nanjing during 17 October 2015 by using the process analysis in WRF-Chem.

According to the reviewer's comments, another two complex air pollution episodes (8-13 July 2015 and 5-11 June 2016) in this region are also selected to conduct simulations for generating general conclusions.

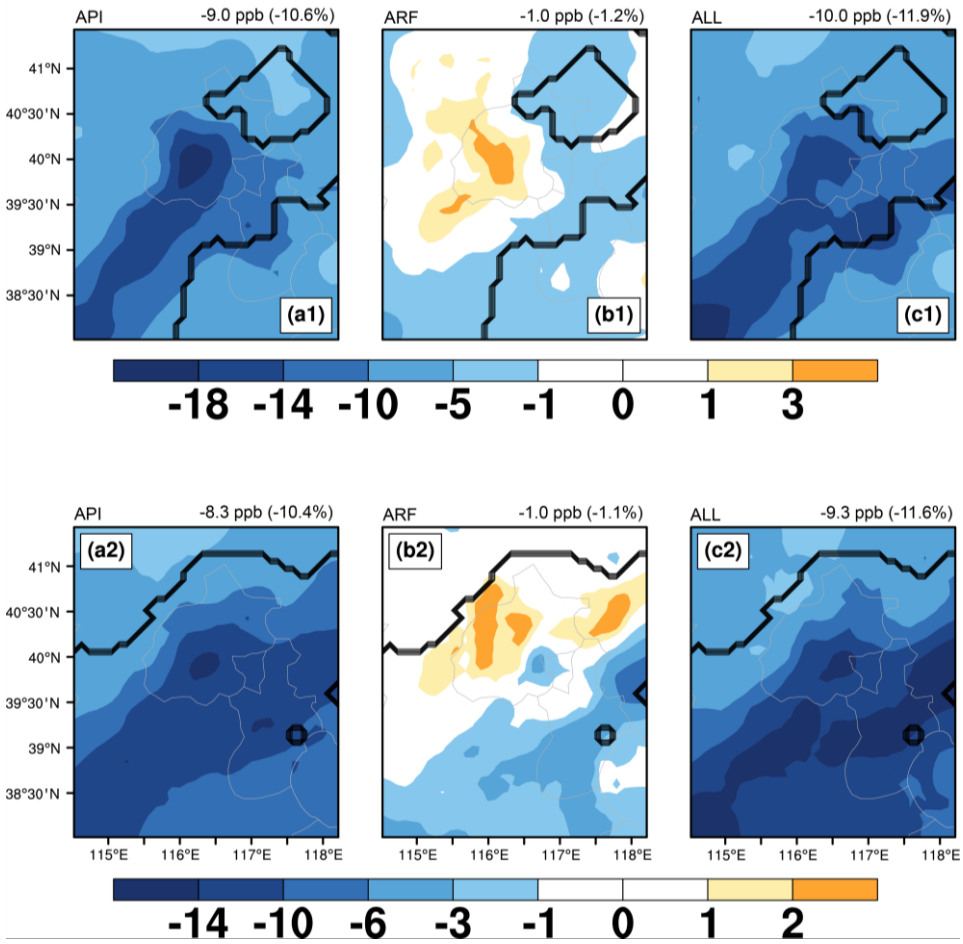

Figure R1. Changes in surface-layer ozone due to (a1-a2) aerosol-photolysis interaction (API), (b1-b2) aerosol-radiation feedback (ARF), and (c1-c2) the combined effects (ALL, defined as API+ARF) during the daytime (08:00-17:00 LST) from 8-13 July 2015 (upper) and 5-11 June 2016 (bottom). The region sandwiched between two black lines is defined as the complex air pollution areas (CAPAs) where the mean daily $PM_{2.5}$ and MDA8 O_3 concentrations in BASE case are larger than 75 μg m^{-3} and 80 ppb. The calculated changes averaged over CAPAs are also shown at the top of each panel.

Simulated air pollutants ($PM_{2.5}$ and O_3) and meteorological variables (T_2, RH_2, and WS_{10}) during 8-13 July 2015 (Episode 2) and 5-11 June 2016 (Episode 3) are compared with observations.

In general, both the observed meteorological parameters and pollutant concentrations can be reasonably reproduced by the model, with correlation coefficients (R) of 0.56~0.98 and normalized mean bias (NMB) of –7.1%~+33.4%. More details about the model evaluation are listed in the supporting information (**Text S1**).

As shown in Fig. R1(a1-a2), API alone leads to the decrease in surface O_3 over the entire domain with an average reduction of 9.0 ppb (10.6%) and 8.3 ppb (10.4%) over CAPAs in Episode 2 and Episode 3, respectively. The decreased surface O_3 concentrations over CAPAs due to ARF are only 1.0 ppb (1.2%, Fig. R1(b1)) and 1.0 ppb (1.1%, Fig. R1(b2)) during Episode 2 and Episode 3, respectively. All the results indicate that API is the dominant factor for O_3 reduction related to aerosol-radiation interactions, the same as the conclusion analyzed from the case during 28 July to 3 August 2014 (Episode 1). The combined effects of API and ARF decrease surface O_3 by 10.0 ppb (11.9%) and 9.3 ppb (11.6%) over CAPAs in Episode 2 and Episode 3, respectively. (**Page 13-14, Line 343-365**)

According to the reviewer's suggestion, Figure R1 is added in the supporting information (**Figure S9**).

2. *Given that government controls have substantially changed emissions in the last decade and will continue into future, how will this research remain relevant in the future or how relevant is it to today's air pollution in China, since the period examined is 7 years ago?*

Response:

The stringent Air Pollution Action Plan has been released by the Chinese government in September 2013 to improve the $PM_{2.5}$ air quality. Although the concentrations of $PM_{2.5}$ are decreasing, the concentrations of $PM_{2.5}$ still exceed 35 μg m^{-3}, and the O_3 levels have continued to increase (Dai et al., 2021). Many studies have found that the decreased $PM_{2.5}$ can be one of the important causes leading to the increase in O_3 (Li et al. 2019; Shao et al., 2021). Li et al. (2019) pointed out that the concentrations of $PM_{2.5}$ were decreased by 40% in North China Plain from 2013 to 2017, which reduced the sink of HO_2 on aerosol surfaces and resulted in the increase in O_3 by analyzing simulation results from the GEOS-Chem model. Meanwhile, the concentrations of O_3 can also be influenced by aerosol-radiation interactions, including aerosol-photolysis interaction and aerosol-radiation feedback, which have not been systematically analyzed. The quantification of the impacts of aerosols on O_3 is important to well understand the co-benefits associated with reductions in both aerosols and O_3.

In this study, we investigate the impacts of aerosol-radiation interactions on surface O_3, and find that the combined impacts of weakened photolysis rates and changed meteorological conditions reduce surface-layer O_3 concentrations by up to 11.4 ppb (13.5%). The result can imply that the decreases in $PM_{2.5}$ can lead to the increase in O_3 due to the weakened aerosol-radiation interactions, which indicates that if the government controls the anthropogenic emissions in future by using the same strategy, higher O_3 will be observed. The result can further emphasize the importance of tighter controls in O_3 precursors (i.e., VOCs) to counteract the increased O_3 caused by weakened aerosol-radiation interactions. Therefore, the contributions of different mitigation strategies with the impacts of aerosol-radiation interactions to O_3 air quality will be discussed detailedly in our future work.

3. *Is the focus of this research the method in which API and ARF are investigated or*

the impact of API and ARF in North China? If it is the former than the authors need to reword the abstract, conclusions, and objectives to make it clear that this study is a "proof-of-concept" study on how to best investigate API and ARF in high O_3 and $PM_{2.5}$ episodes. If the focus is the latter, the authors need to do additional simulations of other high multi-pollutant episodes, perhaps some closer to current conditions and others in the mid 2000s to see if there is change over time or to make the analysis and conclusions more robust.

Response:

This study mainly focuses on the impacts of API and ARF in North China. According to the reviewer's comments, another two complex air pollution episodes (8-13 July 2015 and 5-11 June 2016) in this region are also selected to conduct simulations for generating general conclusions. The impacts of API and ARF on O_3 are shown in Fig. R1, and API is the dominant factor for O_3 reduction related to aerosol-radiation interactions. Similar results can also be concluded by analyzing the episode during 28 July to 3 August 2014.

4. *Does this version of WRF-Chem's CBM-Z and MOSAIC modules have a volatility basis set (VBS) option to simulate secondary organic aerosols and if so is it used? Given that, this is a high O_3 and $PM_{2.5}$ episode there should be a substantial amount of secondary organic aerosol from abundant oxidants and precursors that may be missed in the model without an advanced SOA scheme. How do the author's address the impact of SOA on their conclusions?*

Response:

The selected gas-phase chemical mechanism (CBM-Z) and the aerosol model (MOSAIC) in this study do not consider the impacts of secondary organic aerosols (SOA). The same schemes have been widely used in many other studies, which mainly focus on the impacts of aerosol-radiation interactions on air pollutants in North China (Ding et al., 2016; Gao et al., 2016; Qiu et al., 2017; Chen et al., 2019; Zhou et al, 2019; Gao et al., 2020).

Thanks for the reviewer's suggestion, and we will consider the impacts of SOA in our future works. A discussion about the impacts of SOA has been added in the revised manuscript as follows: "Gao et al. (2017) added some SOA formation mechanisms into the MOSAIC module by using the volatility basis set (VBS) in WRF-Chem and found that the surface $PM_{2.5}$ concentrations in urban Beijing were reduced by 1.9 μg m^{-3} due to the weakened ARF effect during Asia-Pacific Economic Cooperation (APEC). Similar magnitude can also be found in Zhou et al. (2019) (-1.8 μg m^{-3}) who did not consider the impacts of SOA in WRF-Chem when analyzing the impacts of weakened ARF on $PM_{2.5}$ during APEC. Therefore, more work should be conducted to explore the impacts of ARF on $PM_{2.5}$ and O_3 concentrations under consideration of SOA in future." (**Page 15, Line 402-412**)

5. *The authors are investigating aerosol radiation interactions, but the authors do not evaluate the model's performance against either radiation balance datasets or aerosol optical depth. Since these parameters are more important than surface evaluations of air pollutants to understanding API and ARF, the authors should evaluate their model configuration against satellite AOD and radiation variables such as MODIS or CERES-EBAF.*

Response:

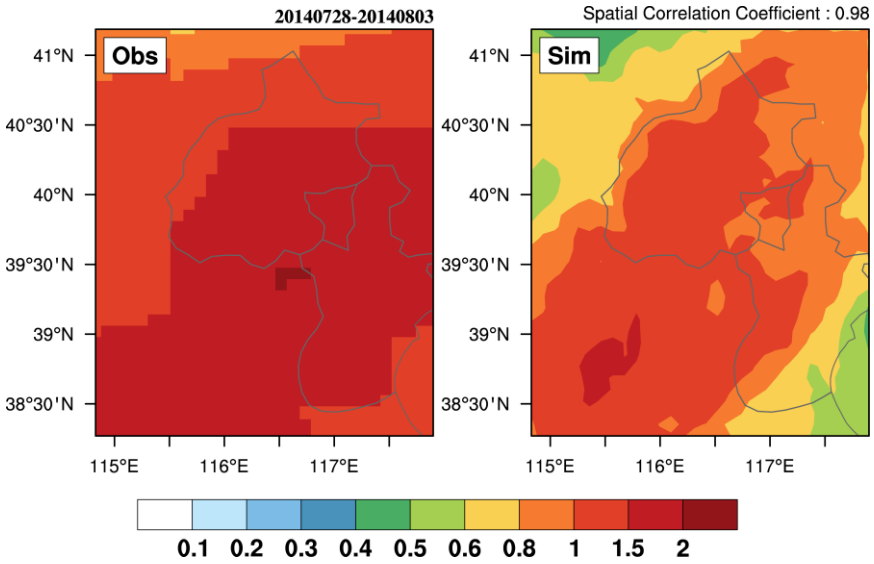

Figure R2. Spatial distributions of aerosol optical depth (AOD) at 550 nm retrieved from MODIS (left) and simulated by WRF-Chem (right) during 28 July to 3 August 2014. The MODIS retrievals are a combination of the standard (over ocean) and "Deep Blue" (over land) products.

Figure R2 shows the spatial distributions of aerosol optical depth (AOD) at 550 nm retrieved from MODIS and simulated by WRF-Chem during 28 July to 3 August 2014. In the WRF-Chem model, the AOD at 550 nm are calculated by using the values at 400 and 600 nm according to the Ångstrom exponent. Analyzing Fig. R2, the model can well reproduce the spatial distribution of observed AOD but slightly underestimate the value. The spatial correlation coefficient between the simulated and observed AOD is 0.98.

According to the reviewer's suggestion, the description of the model evaluation between observed and simulated AOD is added in the revised manuscript (**Page 8, Line 190-196**), and Figure R2 is also added in the supporting information (**Figure S1**).

6. *Are there only three meteorological observation stations in the domain against? If so, why do the authors not also validate their meteorological performance against gridded products like the Climate Research Unit (CRU) datasets to ensure their performance statistics are robust?*

Response:

Thanks to the reviewer's comments. More meteorological observations in the analyzed domain (Table R1) have been used to validate the model results, and the locations of each site are shown in Fig. R3.

Figure R4 shows the time series of observed and simulated T_2, RH_2, WS_{10} and $J[NO_2]$ during 28 July to 3 August 2014. The observed T_2, RH_2 and WS_{10} are averaged from the ten meteorological observation stations, and the photolysis rates of NO_2 are collected from Peking University. Generally, the WRF-Chem model can depict the temporal variation of T_2 fairly well with R of 0.98 and the mean bias (MB) of -1.5 °C. For RH_2, the R and MB are 0.91 and 0.5%, respectively. Although WRF-Chem overestimates WS_{10} with the MB of 0.7 m s^{-1}, the root-mean-square error (RMSE) is 0.9 m s^{-1}, which is smaller than the threshold of model performance criteria (2 m s^{-1}) proposed by Emery et al. (2001). The predicted $J[NO_2]$ agrees well with the observations with R of 0.97 and NMB of 6.8%.

According to the reviewer's comments, we have modified the model evaluation in the revised manuscript (**Page 8, Line 198-208**).

The gridded products like the Climate Research Unit (CRU) datasets covers a large area and a longtime period, which aims to improve scientific understanding of the climate system and its interactions with society. However, the spatial (0.5° × 0.5°) and temporal (monthly) resolution may be too coarse to validate the model performance for generating robust results.

Table R1. Locations of the ten stations from NOAA's National Climatic Data Center used in this study.

Station	Latitude (°)	Longitude (°)
Yuxian	39.833	114.567
Fengning	41.2	116.633
Zhangjiakou	40.783	114.883
Huailai	40.417	115.5
Chengde	40.967	117.917
Beijing	40.08	116.585
Tianjin	39.1	117.167
Binhai	39.124	117.346
Tangshan	39.65	118.1
Baoding	38.733	115.483

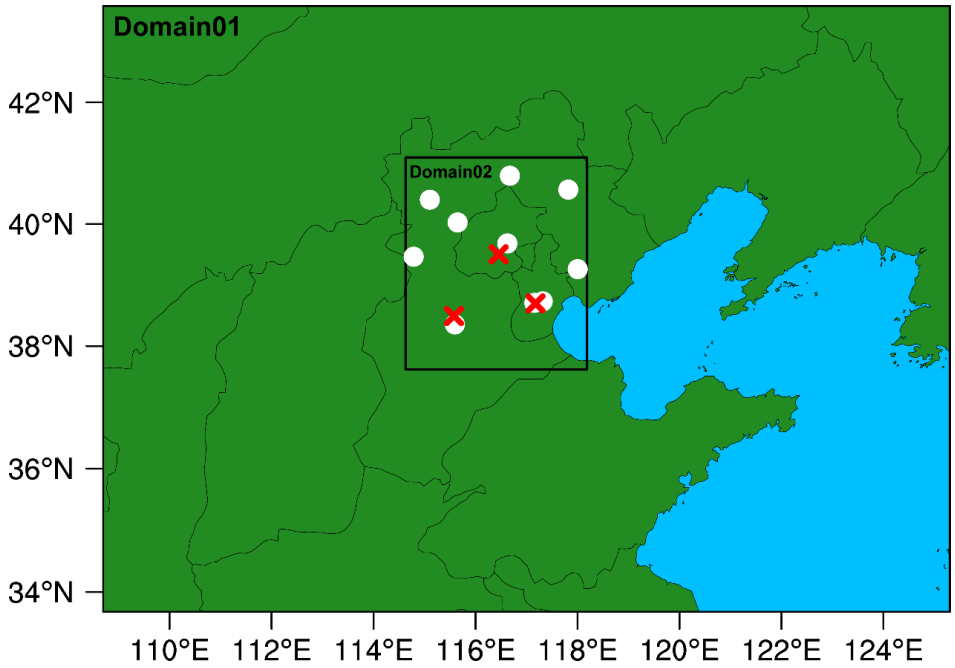

Figure R3. Map of the two WRF-Chem modeling domains with the locations of meteorological (white dots) and environmental (red crosses) observation sites used for model evaluation.

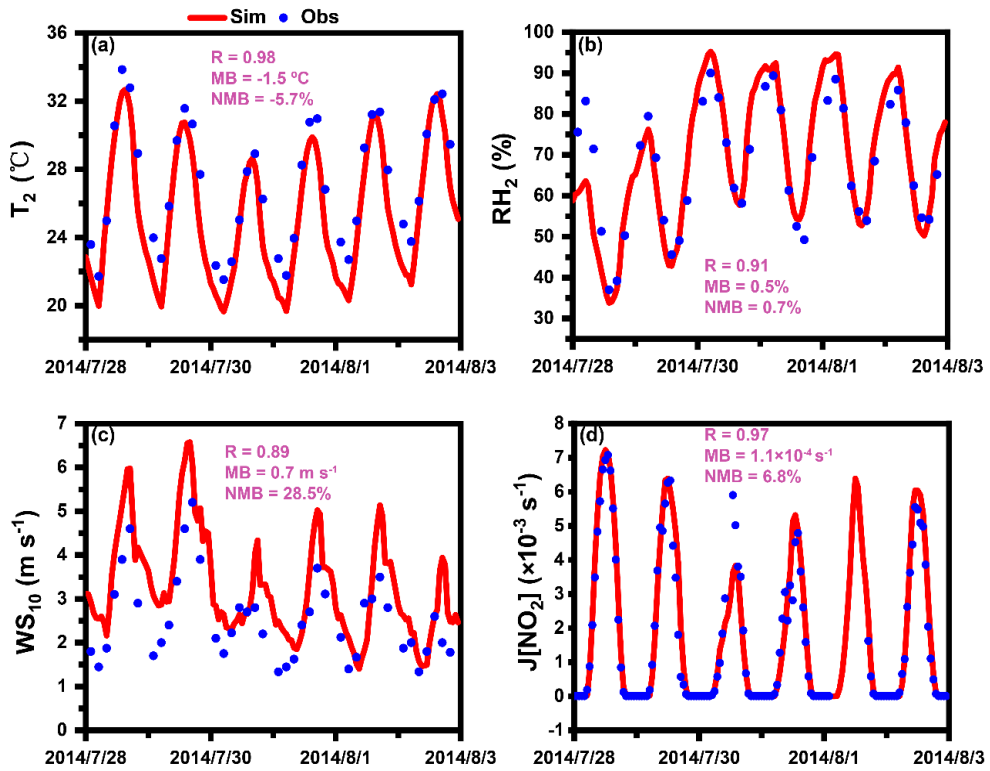

Figure R4. Time series of 3-hourly observed (blue dots) and hourly simulated (red lines) (a) 2-m temperature (T_2), (b) 2-m relative humidity (RH_2), (c) wind speed at 10 m (WS_{10}) averaged over ten meteorological observation stations, and (d) surface photolysis rate of NO_2 ($J[NO_2]$) during 28 July to 3 August 2014. The calculated correlation coefficient (R), mean bias (MB), and normalized mean bias (NMB) are also shown.

7. *Given that interactions between O_3 and $PM_{2.5}$ are non-linear, how do the authors justify using a simple ratio value (i.e., ROP) to relate these interactions? If this ratio does not account for non-linearity, how useful is this value?*

Response:

Thanks to the reviewer's suggestion. As the relationship between O_3 and $PM_{2.5}$ is non-linear, and the simple index of ROP can not fully represent the impacts of aerosols on surface O_3, so we delete the ROP in the revised manuscript.

8. *The axis labels and legends of Figure 7 are difficult to read. Either each panel should be larger overall or the font sizes of the axes and legends need increased.*

Response:

According to the reviewer's suggestion, we have modified the axis labels and legends of Figure 7 and the other figures in the revised manuscript. (**Page 32**)

Minor Comments:

1) *In the abstract, there is no context for the values listed. Further reading into the manuscript reveals that these values are the averages in the areas of the complex air pollution areas. The authors should briefly state that these values are for daytime average changes in complex air pollution areas in the abstract. I would also suggest adding a more processed based explanation of the*

changes in atmospheric state rather than simply listing a long series of values. For example, the authors could state something similar to the following: "Aerosol radiation interactions lead to shortwave dimming at the earth's surface of X, which reduce photolysis rates by X. The dimming stabilizes the atmosphere via surface cooling of X, which reduces PBL height by X. The stabilized atmosphere increases saturation in the lower atmosphere by X. etc…."

Response:

According to the reviewer's suggestion, we have revised the explanation in the abstract as follows: "Our results show that aerosol-radiation interactions decreased the daytime shortwave radiation at surface by 93.2 W m^{-2} averaged over the complex air pollution areas. The dimming effect reduced the 2 m temperature and near-surface photolysis rates of J[NO_2] and J[O^1D] by 0.56 °C, 1.8 × 10^{-3} s^{-1} and 6.1 × 10^{-6} s^{-1}, respectively. However, the daytime shortwave radiation in the atmosphere was increased by 72.8 W m^{-2}, which made the atmosphere more stable. The stabilized atmosphere decreased the planetary boundary layer height and 10 m wind speed by 129.0 m and 0.12 m s^{-1}, respectively, and increased the relative humidity at 2 m by 2.4%." (**Page 2, Line 24-32**)

2) *Make it clear throughout the manuscript when you are referring to surface level O_3 and $PM_{2.5}$.*

Response:

According to the reviewer's suggestion, we have revised the expressions in the whole manuscript.

3) *Lines 179-181: The missing $PM_{2.5}$ could also be from missing SOA formation pathways, as mentioned above, if no advanced SOA formulations are used.*

Response:

Thanks for your suggestion. The selected aerosol model (MOSAIC) in this study does not consider the impacts of secondary organic aerosols (SOA), and we have added the explanation in the revised manuscript as follows: "The failure to reproduce $PM_{2.5}$ peak values may be attributed to incomplete treatments of chemical reactions in WRF-Chem, e.g., missing the heterogeneous chemistry in the model (Cheng et al., 2016) and the lack of secondary organic aerosols (SOA) formation pathways in the aerosol module (Chen et al., 2019)." (**Page 7-8, Line 183-188**)

4) *Is "downward shortwave radiation in the atmosphere" the SWDNT variable from WRF-Chem? If so, the name of this variable is "downward shortwave radiation at the top of the atmosphere".*

Response:

Thanks for your comments. In the WRF-Chem model, SWDNT (SWUPT) means the download (upward) shortwave radiation at the top of atmosphere, and SWDNB (SWUPB) represents the download (upward) shortwave radiation at the surface. According to Zhao et al. (2011), the shortwave radiation in the atmosphere (ATM_SW) can be calculated as the difference between TOP_SW (the net shortwave radiation at the top of atmosphere, i.e.,

SWDNT minus SWUPT) and BOT_SW (the net shortwave radiation at the surface, i.e., SWDNB minus SWUPB).

According to the reviewer's suggestion, we have changed the expressions of BOT_SW (shortwave radiation at the surface) and ATM_SW (shortwave radiation in the atmosphere) in the whole revised manuscript.

5) *Lines 217-218: If ATM_SW is the SWDNT variable, what is causing it to increase? SWDNT is usually controlled by the solar constant. Is it possible this is reflected upward shortwave (SWUPT)?*

Response:

ATM_SW represents the shortwave radiation in the atmosphere, and it can be calculated by the following equation: ATM_SW = (SWDNT - SWUPT) – (SWDNB - SWUPB).

6) *Lines 248-249: This should be revised to make it clearer that ARF primarily impacts O_3 through changing the NOx distribution.*

Response:

According to the comments of Reviewer#1, we have deleted this sentence.

7) *Lines 270-281: Is VMIX increasing surface O_3 because it is mixing down higher O_3 concentrations from aloft or because vertical mixing is suppressed due to a stable atmosphere?*

Response:

VMIX increases the surface O_3 concentrations by transporting the higher O_3 from aloft to the surface layer. Similar results can also be found in previous studies (Tang et al., 2017; Xing et al., 2017; Gao et al., 2018).

8) *Lines 282-294: Why does the VMIX contribution increase because of API?*

Response:

Analyzing the vertical profiles of the differences in contributions from each physical/chemical process to hourly O_3 variations caused by API in Fig. 8(b), we found that the contribution of VMIX_DIF is negative in the aloft (among the 9th and the 13th layers), while it turns to be positive at the lower seven layers, and the positive contribution increases as the height decreases. The positive variation in VMIX due to API may be associated with the different vertical gradient of O_3 between BASE and NOAPI cases.

Similar results can also be found in Gao et al. (2020), who concluded that the increased vertical gradients of O_3 due to API could enhance the vertical entrainment.

9) *Lines 295-301: Explain why VMIX_DIF and CHEM_DIF are positive during the day due to ARF.*

Response:

When the impacts of ARF are considered, PBLH is decreased over CAPAs (Fig. S3(b3)), which indicates that the suppressed PBL in NOAPI restrains the vertical turbulence and prevents O_3 being transported from aloft to surface, resulting in lower O_3 concentrations at surface when comparing with the simulation results of NOALL. However, as the evolution in boundary layer

during the daytime, more O_3 can be diffused from the upper layers to the surface in NOAPI, and the differences in hourly variation in surface O_3 due to vertical mixing between NOAPI and NOALL are positive. Similar results can also be found in Gao et al. (2018).

The typical VOCs/NO_x ratio is calculated to classify sensitivity regimes and to indicate the possible O_3 responses to changes in VOCs and/or NO_x concentrations. O_3 production is VOC-limited if the ratio is less than 4, and it is NO_x-limited if the ratio is larger than 15 (Edson et al., 2017; Li et al., 2017). The ratio of VOCs/NO_x ranging around 4-15 indicates a transitional regime, where ozone is nearly equally sensitive to each species (Sillman, 1999). As shown in Fig R5(a-c), O_3 are mainly formed under the VOC-limited and the transition regimes in CAPAs, which means that the increased concentrations of VOCs and NO_x are favorable for ozone chemical production. As shown in Fig. R5(e) and (h), both the surface concentrations of VOCs and NO_x are increased when the impacts of ARF are considered. Thus, the contribution of CHEM in NOAPI is larger than that in NOALL. Similar results can also be found in Gao et al. (2018).

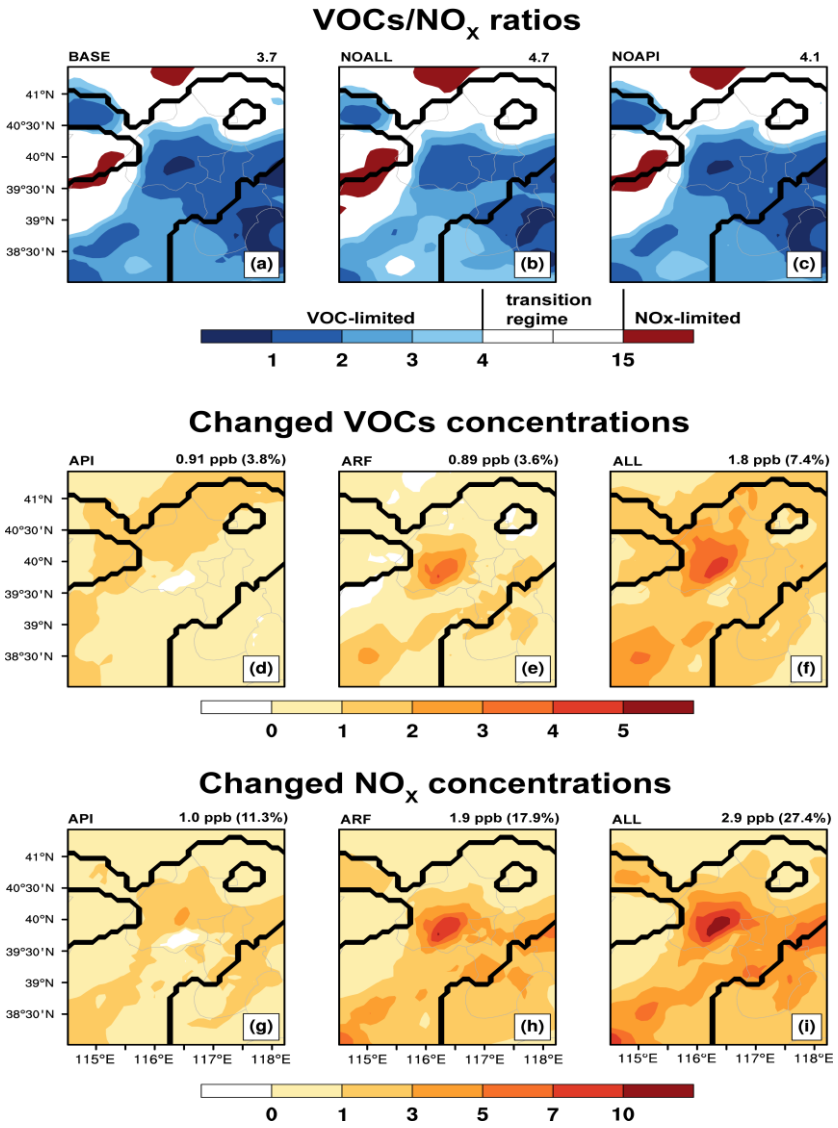

Figure R5. The ratios of VOCs/NO_x calculated from (a) BASE, (b) NOALL, and (c) NOAPI. The changed surface-layer concentrations of VOCs and NO_x (NO_2+NO, ppb) caused by (d, g) API, (e, h) ARF, and (f, i) ALL during the

daytime (08:00-17:00 LST) from 28 July to 3 August 2014. The calculated values averaged over CAPAs are also shown at the top of each panel.

10) *Lines 315-316: Explain how different vertical O_3 gradients can cause this change.*

Response:

Since the VMIX is closely dependent on atmospheric turbulence and vertical gradients of O_3 concentration. The API will increase vertical gradients of O_3 to enhance the vertical entrainment (Gao et al., 2020).

Line Comments:

1) *Line 49: This should be "Earth's radiative balance" or "Earth's energy balance"*

Response:

Thanks for your suggestion. We have changed the expression in the revised manuscript. (**Page 3, Line 48**)

2) *Lines 54-56: Are these studies all focused on China? If so, state that in the sentence. Change "were" to "are".*

Response:

According to the reviewer's suggestion, we have changed the expression in the revised manuscript. (**Page 3, Line 54**)

3) *Lines 56-63: State the domain and time period of Gao et al., (2015) at the beginning of this statement rather than the end*

Response:

According to the reviewer's suggestion, we have changed the expression in the revised manuscript. (**Page 3, Line 55-62**)

4) *Line 66: Add "the" before North China Plain*

Response:

Thanks for your suggestion. We have added the "the" before North China Plain in the revised manuscript. (**Page 3, Line 64**)

5) *Lines 66-67: If this is referring to surface $PM_{2.5}$ concentrations, add "surface" before $PM_{2.5}$ concentrations.*

Response:

Thanks for your suggestion. We have added the "surface" before $PM_{2.5}$ concentrations in the revised manuscript. (**Page 3, Line 65**)

6) *Line 204: should be "attention"*

Response:

Thanks for your suggestion. We have changed the expression in the revised manuscript. (**Page 9, Line 218**)

7) *Line 256: Center align the equation.*

Response:

This equation has been deleted.

8) *Line: 259: Why are there parentheses in the units?*

Response:

This sentence has been deleted.

9) *Lines 288-289: This sentence is a little confusing. Is Net_DIF the sum of CHEM_DIF, VMIX_DIF, and ADV_DIF? If so, state that explicitly and then indicate what Net_DIF describes.*

Response:

Thanks for your suggestion. We have defined the NET_DIF in the revised manuscript. (**Page 11, Line 291-292**)

10) *Line 321: Remove "in the"*

Response:

According to the reviewer's suggestion, we have deleted it in the revised manuscript.

11) *Line 361: Remove "the contribution from VMIX and"*

Response:

According to the reviewer's suggestion, we have deleted it in the revised manuscript.

12) *Line 373: Either "A recent study" or "Recent studies have"*

Response:

According to the reviewer's suggestion, we have changed the expression in the revised manuscript. (**Page 15, Line 396**)

Reference:

Chen, L., Zhu, J., Liao, H., Gao, Y., Qiu, Y., Zhang, M., Liu, Z., Li, N., and Wang, Y.: Assessing the formation and evolution mechanisms of severe haze pollution in the Beijing–Tianjin–Hebei region using process analysis, Atmos. Chem. Phys., 19, 10845–10864, https://doi.org/10.5194/acp-19-10845-2019, 2019.

Cheng, Y., Zheng, G., Chao, W., Mu, Q., Bo, Z., Wang, Z., Meng, G., Qiang, Z., He, K., and Carmichael, G.: Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events in China, Science Advances, 2, https://doi.org/10.1126/sciadv.1601530, 2016.

Dai, H., Zhu, J., Liao, H., Li, J., Liang, M., Yang, Y., and Yue, X.: Co-occurrence of ozone and $PM_{2.5}$ pollution in the Yangtze River Delta over 2013–2019: Spatiotemporal distribution and meteorological conditions, Atmos. Res., 249, 105363, 2021.

Ding, A. J., Huang, X., Nie,W., Sun, J. N., Kerminen, V. M., Petäjä, T., Su, H., Cheng, Y. F., Yang, X. Q., Wang, M. H., Chi, X. G., Wang, J. P., Virkkula, A., Guo, W. D., Yuan, J., Wang, S. Y., Zhang, R. J., Wu, Y. F., Song, Y.,

Zhu, T., Zilitinkevich, S., Kulmala, M., and Fu, C. B.: Enhanced haze pollution by black carbon in megacities in China, Geophys. Res. Lett., 43, 2873–2879, https://doi.org/10.1002/2016gl067745, 2016.

Edson, C. T., Ivan, H.-P. and Alberto, M.: Use of combined observational- and model-derived photochemical indicators to assess the O_3-NO_x-VOC System sensitivity in urban areas, Atmosphere., 8, 22. https://doi.org/10.3390/ atmos8020022, 2017.

Emery, C., Tai, E., and Yarwood, G.: Enhanced meteorological modeling and performance evaluation for two Texas ozone episodes, in: Prepared for the Texas Natural Resource Conservation Commission, ENVIRON International Corporation, Novato, CA, USA, 2001.

Gao, J. H., Zhu, B., Xiao, H., Kang, H. Q., Pan, C., Wang, D. D., and Wang, H. L.: Effects of black carbon and boundary layer interaction on surface ozone in Nanjing, China, Atmos. Chem. Phys., 18, 7081–7094, https://doi.org/10.5194/acp-18-7081-2018, 2018.

Gao, J., Li, Y., Zhu, B., Hu, B., Wang, L., and Bao, F.: What have we missed when studying the impact of aerosols on surface ozone via changing photolysis rates?, Atmos. Chem. Phys., 20, 10831–10844, https://doi.org/10.5194/acp-20-10831-2020, 2020.

Gao, M., Carmichael, G. R., Wang, Y., Saide, P. E., Yu, M., Xin, J., Liu, Z., and Wang, Z.: Modeling study of the 2010 regional haze event in the North China Plain, Atmos. Chem. Phys., 16, 1673–1691, doi:10.5194/acp-16-1673-2016, 2016.

Gao, M., Liu, Z., Wang, Y., Lu, X., Ji, D. and Wang, L.: Distinguishing the roles of meteorology, emission control measures, regional transport, and co-benefits of reduced aerosol feedbacks in “APEC” Blue, Atmos. Environ., 167, 476–486, 420 doi:10.1016/j.atmosenv.2017.08.054, 2017.

Li, K., Chen, L., Ying, F., White, S. J., Jang, C., Wu, X., Gao, X., Hong, S., Shen, J., Azzi, M. and Cen, K: Meteorological and chemical impacts on ozone formation: a case study in Hangzhou, China, Atmos. Res., 196, https://doi.org/10.1016/ j.atmosres.2017.06.003, 2017.

Li, K., Jacob, D. J., Liao, H., Shen, L., Zhang, Q., and Bates, K. H.: Anthropogenic drivers of 2013–2017 trends in summer surface ozone in China, P. Natl. Acad. Sci. USA, 116, 422–427, https://doi.org/10.1073/pnas.1812168116, 2019.

Qiu, Y., Liao, H., Zhang, R., and Hu, J.: Simulated impacts of direct radiative effects of scattering and absorbing aerosols on surface layer aerosol concentrations in China during a heavily polluted event in February 2014, J. Geophys. Res. Atmos., 122, 5955–5975, doi:10.1002/2016JD026309, 2017.

Shao, M., Wang, W. J., Yuan, B., Parrish, D. D., Li, X., Lu, K. D., Wu, L. L., Wang, X. M., Mo, Z. W., Yang, S. X., Peng, Y. W., Kuang, Y., Chen, W. H., Hu, M., Zeng, L. M., Su, H., Cheng, Y. F., Zheng, J. Y., Zhang, Y. H.: Quantifying the role of $PM_{2.5}$ dropping in variations of ground-level ozone: Inter-comparison between Beijing and Los Angeles, Sci. Total Environ., https://doi.org/10.1016/j.scitotenv.2021.147712, 2021.

Sillman, S.: The relation between ozone, NO_x and hydrocarbons in urban and polluted rural environments, Atmos. Environ., 33, 1821-1845, https://doi.org/ 10.1016/S1352-2310(98)00345-8, 1999.

Tang, G. Q., Zhu, X.W., Xin, J. Y., Hu, B., Song, T., Sun, Y., Zhang, J. Q., Wang, L. L., Cheng, M. T., Chao, N., Kong, L. B., Li, X., and Wang, Y. S.: Modelling study of boundary-layer ozone over northern China – Part I: Ozone budget in summer, Atmos. Res., 187, 128–137, 2017.

Wang, T., Huang, X., Wang, Z., Liu, Y., Zhou, D., Ding, K., Wang, H., Qi, X. and Ding, A.: Secondary Aerosol Formation and its Linkage with Synoptic Conditions during Winter Haze Pollution over Eastern China, Sci. Total Environ., 730, 138888, 2020.

Xing, J., Wang, J. D., Mathur, R., Wang, S. X., Sarwar, G., Pleim, J., Hogrefe, C., Zhang, Y. Q., Jiang, J. K., Wong, D. and Hao, J. M.: Impacts of aerosol direct effects on tropospheric ozone through changes in atmospheric

dynamics and photolysis rates, Atmos. Chem. Phys., 17, 9869–9883, https://doi.org/10.5194/acp-17-9869-2017, 2017.

Zhao, C., Liu, X., Ruby Leung, L., and Hagos, S.: Radiative impact of mineral dust on monsoon precipitation variability over West Africa, Atmos. Chem. Phys., 11, 1879–1893, https://doi.org/10.5194/acp-11-1879-2011, 2011.

Zhou, M., Zhang, L., Chen, D., Gu, Y., Fu, T.-M., Gao, M., Zhao, Y., Lu, X. and Zhao, B.: The impact of aerosol-radiation interactions on the effectiveness of emission control measures, Environmental Research Letters, 14(2), 024002, https://doi.org/10.1088/1748-9326/aaf27d, 2019.

Thank you very much for your comments and suggestions.

Impacts of aerosol-photolysis interaction and aerosol-radiation feedback on surface-layer ozone in North China during a multi-pollutant air pollution episode

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Abstract

We examined the impacts of aerosol-radiation interactions, including the effects of aerosol-photolysis interaction (API) and aerosol-radiation feedback (ARF), on surface-layer ozone (O_3) concentrations during one multi-pollutant air pollution episode characterized by high O_3 and $PM_{2.5}$ levels from 28 July to 3 August 2014 in North China, by using the Weather Research and Forecasting with Chemistry (WRF-Chem) model embedded with an integrated process analysis scheme. Our results show that aerosol-radiation interactions decreased the daytime shortwave radiation at surface by 93.2 W m^{-2} averaged over the complex air pollution areas. The dimming effect reduced the 2 m temperature and near-surface photolysis rates of $J[NO_2]$ and $J[O^1D]$ by 0.56 °C, 1.8 × 10^{-3} s^{-1} and 6.1 × 10^{-6} s^{-1}, respectively. However, the daytime shortwave radiation in the atmosphere was increased by 72.8 W m^{-2}, which made the atmosphere more stable. The stabilized atmosphere decreased the planetary boundary layer height and 10 m wind speed by 129.0 m and 0.12 m s^{-1}, respectively, and increased the relative humidity at 2 m by 2.4%. ~~Our results show that aerosol-radiation interactions decrease the daytime downward shortwave radiation at surface, 2 m temperature, 10 m wind speed, planetary boundary layer height, photolysis rates $J[NO_2]$ and $J[O^1D]$ by 115.8 W m^{-2}, 0.56 °C, 0.12 m s^{-1}, 129 m, 1.8 × 10^{-3} s^{-1} and 6.1 × 10^{-6} s^{-1}, and increase relative humidity at 2 m and downward shortwave radiation in the atmosphere by 2.4% and 72.8 W m^{-2}.~~ The weakened photolysis rates and changed meteorological conditions reduced daytime surface-layer O_3 concentrations by up to 11.4 ppb (13.5%), with API and ARF contributing 74.6% and 25.4% of the O_3 decrease, respectively. ~~The combined impacts of API and ARF on surface O_3 are further quantitatively characterized by the ratio of changed O_3 concentration to local $PM_{2.5}$ level. The ratio is calculated to be -0.14 ppb (μg m^{-3})$^{-1}$ averaged over the multi-pollutant air pollution area in North China.~~ Process analysis ~~indicates~~ indicated that the weakened O_3 chemical production ~~makes~~ made the greatest contribution to API effect while the reduced vertical mixing ~~is~~ was the key process for ARF effect. This study implies that future $PM_{2.5}$ reductions will lead to O_3 increases due to weakened aerosol-radiation interactions. Therefore, tighter controls of

47 O_3 precursors are needed to offset O_3 increases caused by weakened aerosol-radiation
48 interactions in the future.

1 Introduction

China has been experiencing severe air pollution in recent years, characterized by high loads of $PM_{2.5}$ (particulate matter with an aerodynamic equivalent diameter of 2.5 micrometers or less) and high levels of ozone (O_3). Observational studies exhibited positive correlations and synchronous occurrence of $PM_{2.5}$ and O_3 pollution in North China during summer (Zhao et al., 2018; Zhu et al., 2019), indicating that complex air pollution is becoming a major challenge for North China.

Aerosols can absorb and scatter solar radiation ~~and therefore alter~~to affect Earth's energy balance~~radiative balance~~. They can also act as cloud condensation nuclei and ice nuclei, and further modify the microphysical characteristics of clouds (Albrecht et al., 1989; Haywood et al., 2000; Lohmann et al., 2005). Both ways perturb meteorological variables, e.g., temperature, planetary boundary layer height (PBLH), and precipitation, and eventually influence air pollutants (Petäjä et al., 2015; Miao et al., 2018; Zhang et al., 2018). Many studies ~~were~~ are focused on the feedback between aerosol and meteorology (Gao et al., 2015; Gao et al., 2016a; Qiu et al., 2017; Chen et al., 2019; Zhu et al., 2021). Gao et al. (2015) used the WRF-Chem model to investigate the feedbacks between aerosols and meteorological variables over the North China Plain in January 2013, and pointed out that aerosols could cause a decrease in surface temperature by 0.8-2.8 °C but an increase of 0.1-0.5 °C around 925 hPa ~~when feedbacks between aerosols and meteorological variables were considered in WRF-Chem model~~. The more stable atmosphere caused by surface cooling and higher-layer heating led to the decreases of surface wind speed and PBLH by 0.3 m s^{-1} and 40-200 m, respectively, which further resulted in overall $PM_{2.5}$ increases by 10-50 μg m^{-3} (2-30%) ~~over Beijing, Tianjin and south Hebei during January 2013~~. By using the same WRF-Chem model, Qiu et al. (2017) reported that the surface downward shortwave radiation and PBLH were reduced by 54.6 W m^{-2} and 111.4 m due to aerosol radiative forcing during 21 and 27 February 2014 in the North China Plain. As a result, the surface $PM_{2.5}$ concentration averaged over the North China Plain was increased by 34.9 μg m^{-3} (20.4%).

Aerosols can also influence O_3 through aerosol-radiation interactions, including

aerosol-photolysis interaction and aerosol-radiation feedback. Aerosols can scatter and absorb UV radiation, and therefore directly affect O_3 photochemistry reactions, which is called aerosol-photolysis interaction (API) (Dickerson et al., 1997; Liao et al., 1999; Li et al., 2011; Lou et al., 2014). The changed meteorological variables due to aerosol radiative forcing can indirectly affect O_3 concentrations, which is called aerosol-radiation feedback (ARF) (Hansen et al., 1997; Gao et al., 2018; Liu et al., 2020). Although the effects of API or ARF on O_3 have been examined by previous studies (Xing et al., 2017; Gao et al., 2018; Gao et al., 2020), the combined effects of API and ARF on O_3, especially under the conditions of synchronous occurrence of high $PM_{2.5}$ and O_3 concentrations, remain largely elusive.

The present study aims to (1) quantify the respective/combined contributions of API and ARF on surface O_3 concentrations by using the WRF-Chem model; (2) explore the prominent physical and/or chemical processes responsible for API and ARF effects by using an integrated process rate (IPR) analysis embedded in WRF-Chem model. The analysis is conducted during one multi-pollutant air pollution episode characterized by high O_3 and $PM_{2.5}$ levels from 28 July to 3 August 2014 in North China. The model configuration, numerical experiments, observational data, and the integrated process rate analysis are described in section 2. Section 3 shows the model evaluation. ~~The presentation and discussion of the model results are exhibited in section 4, and the conclusion is provided in section 5.~~ The presentation of the model results and the corresponding analyses are exhibited in section 4. The discussion is provided in section 5, and the conclusion and uncertainties of this study are given in section 6.

2 Methods

2.1 Model configuration

The version 3.7.1 of the online-coupled Weather Research and Forecasting with Chemistry (WRF-Chem) model (Grell et al., 2005; Skamarock et al., 2008) is used in this study to explore the impacts of aerosol-radiation interactions on surface-layer O_3 in North China. WRF-Chem can simulate gas phase species and aerosols coupled with meteorological fields, and has been widely used to investigate air pollution over North

China (Gao et al., 2016a; Gao et al., 2020; Wu et al., 2020). As shown in Fig. 1, we design two nested model domains with the number of grid points of 57 (west–east) × 41 (south–north) and 37 (west–east) × 43 (south–north) at 27 and 9 km horizontal resolutions, respectively. The parent domain centers at 39 °N, 117 °E. The model contains 29 vertical levels from the surface to 50 hPa, with 14 levels below 2 km for the fully description of the vertical structure of planetary boundary layer (PBL).

The Carbon Bond Mechanism Z (CBM-Z) is selected as the gas-phase chemical mechanism (Zaveri and Peters, 1999), and the full 8-bin MOSAIC (Model for Simulating Aerosol Interactions and Chemistry) aerosol module with aqueous chemistry is used to simulate aerosol evolution (Zaveri et al., 2008). The photolysis rates are calculated by the Fast-J scheme (Wild et al., 2000). Other major physical parameterizations used in this study are listed in Table 1.

The initial and boundary meteorological conditions are provided by the National Centers for Environmental Prediction (NCEP) Final Analysis data with a spatial resolution of 1° × 1°. In order to limit the model bias of simulated meteorological fields, the four-dimensional data assimilation (FDDA) is used with a nudging coefficient of 3.0×10^{-4} for the wind, temperature and humidity (no analysis nudging is applied for the inner domain) (Lo et al., 2008; Otte, 2008). Chemical initial and boundary conditions are obtained from the Model for Ozone and Related chemical Tracers, version 4 (MOZART-4) forecasts (Emmons et al., 2010).

Anthropogenic emissions are taken from the 2010 MIX Asian emission inventory (Li et al., 2017a), which provides emissions of sulfur dioxide (SO_2), nitrogen oxides (NO_x), carbon monoxide (CO), non-methane volatile organic compounds (NMVOCs), carbon dioxide (CO_2), ammonia (NH_3), black carbon (BC), organic carbon (OC), PM_{10} (particulate matter with aerodynamic diameter is 10 μm and less) and $PM_{2.5}$. Emissions are aggregated from four sectors, including power generation, industry, residential, and transportation, with 0.25° × 0.25° spatial resolution. Biogenic emissions are calculated online by the Model of Emissions of Gases and Aerosols from Nature (MEGAN) (Guenther et al., 2006).

2.2 Numerical experiments

To quantify the impacts of API and ARF on O_3, three case simulations have been conducted: (1) BASE – the base simulation coupled with the interactions between aerosol and radiation, which includes both impacts of API and ARF; (2) NOAPI – the same as the BASE case, but the impact of API is turned off (aerosol optical properties are set to zero in the photolysis module), following Wu et al. (2020); (3) NOALL – both the impacts of API and ARF are turned off (removing the mass of aerosol species when calculating aerosol optical properties in the optical module), following Qiu et al. (2017). The differences between BASE and NOAPI (i.e., BASE minus NOAPI) represent the impacts of API. The contributions from ARF can be obtained by comparing NOAPI and NOALL (i.e., NOAPI minus NOALL). The combined effects of API and ARF on O_3 concentrations can be quantitatively evaluated by the differences between BASE and NOALL (i.e., BASE minus NOALL).

Each simulation is conducted from 26 July to 3 August 2014, with the first 40 hours as the model spin-up. Simulation results from the BASE case during 28 July and 3 August 2014 are used to evaluate the model performance.

2.3 Observational data

Simulation results are compared with meteorological and chemical measurements. The surface-layer meteorological data (2 m temperature (T_2), 2 m relative humidity (RH_2), and 10 m wind speed (WS_{10})), with a temporal resolution of 3 h, at ~~three~~ ten stations (Table S1) are obtained from NOAA's National Climatic Data Center (https://gis.ncdc.noaa.gov/maps/ncei/cdo/hourly). The radiosonde data of temperature at 08:00 and 20:00 LST in Beijing (39.93 °N, 116.28 °E) are provided by the University of Wyoming (http://weather.uwyo.edu/). Observed hourly concentrations of $PM_{2.5}$ and O_3 at thirty-two sites (Table S2) in North China are collected from the China National Environmental Monitoring Center (CNEMC). The photolysis rate of nitrogen dioxide (NO_2) ($J[NO_2]$) measured at the Peking University site (39.99 °N, 116.31 °E) is also used to evaluate the model performance. More details about the measurement technique of $J[NO_2]$ can be found in Wang et al. (2019). The satellite-retrieved 550 nm AOD

products from the Moderate Resolution Imaging Spectroradiometer (MODIS) are also used to compare with the simulated ones. The model results from 10:00 to 11:00 and 13:00 to 14:00 LT are extracted and averaged, due to instruments on board the Terra and Aqua platforms pass over China at around 10:30 and 13:30 LT, respectively.

2.4 Integrated process rate analysis

Integrated process rate (IPR) analysis has been widely used to quantify the contributions of different processes to O_3 variations (Goncalves et al., 2009; Gao et al., 2016b; Tang et al., 2017; Gao et al., 2018). In this study, four physical/chemical processes are considered, including vertical mixing (VMIX), net chemical production (CHEM), horizontal advection (ADVH), and vertical advection (ADVZ). VMIX is initiated by turbulent process and closely related to PBL development, which influences O_3 vertical gradients. CHEM represents the net O_3 chemical production (chemical production minus chemical consumption). ADVH and ADVZ represent transport by winds (Gao et al., 2016b). In this study, we define ADV as the sum of ADVH and ADVZ.

3 Model evaluation

Reasonable representation of observed meteorological and chemical variables by the WRF-Chem model can provide foundation for evaluating the impacts of aerosols on surface-layer ozone concentration. The model results presented in this section are taken from the BASE case. The concentrations of air pollutants are averaged over the thirty-two observation sites in Beijing, Tianjin and Baoding. To ensure the data quality, the mean value for each time is calculated only when concentrations are available at more than sixteen sites.

3.1 Chemical simulations

Figure 2 shows the spatial-temporal variations of observed and simulated $PM_{2.5}$ and O_3 concentrations over North China during 28 July to 3 August 2014. The observed higher concentrations in Beijing and Baoding than those in Tianjin are well reproduced by ~~the WRF-Chem model~~WRF-Chem. The model can also reasonably capture the temporal variations of observed $PM_{2.5}$ and O_3 with ~~high~~ correlation coefficients (R) of

0.66 for $PM_{2.5}$ and 0.86 for O_3, although simulated results underestimate the observed $PM_{2.5}$ by -19.2% and O_3 by -12.0%. The failure to reproduce $PM_{2.5}$ peak values may be attributed to incomplete treatments of chemical reactions in WRF-Chem, e.g., The failure to reproduce $PM_{2.5}$ peak values may be attributed to incomplete treatments of chemical reactions in WRF-Chem, e.g., missing the heterogeneous chemistry in the model (Cheng et al., 2016) and the lack of secondary organic aerosols (SOA) formation pathways in the aerosol module (Chen et al., 2019).~~the aqueous-phase reactions of SO_2 oxidized by NO_2 in aerosol water (Cheng et al., 2016).~~ More statistical parameters between simulations and observations are presented in Table 2.

Figure S1 shows the spatial distributions of aerosol optical depth (AOD) at 550 nm retrieved from MODIS and simulated by WRF-Chem during 28 July to 3 August 2014. In the WRF-Chem model, the AOD at 550 nm are calculated by using the values at 400 and 600 nm according to the Ångstrom exponent. Analyzing Fig. S1, the model can well reproduce the spatial distribution of observed AOD but slightly underestimate the value. The spatial correlation coefficient between the simulated and observed AOD is 0.98.

3.2 Meteorological simulations

Figure 3 shows the time series of observed and simulated T_2, RH_2, and WS_{10} averaged~~ over three cities (Beijing, Tianjin, and Baoding)~~ over ten meteorological observation stations, and J[NO_2] at Peking University during 28 July to 3 August 2014. The statistical metrics for T_2, RH_2, WS_{10}, and J[NO_2] are also presented in Table 2. Generally, the model can depict the temporal variations of T_2 fairly well with R of 0.98 and the mean bias (MB) of ~~-0.2~~-1.5 °C. For RH_2, the R and MB are ~~0.93~~0.91 and ~~-6.00~~0.5%, respectively. Although WRF-Chem model overestimates WS_{10} with the MB of ~~0.6~~0.7 m s^{-1}, the R for WS_{10} is ~~0.70~~0.89 and the root-mean-square error (RMSE) is ~~1.00~~0.9 m s^{-1}, which is smaller than the threshold of model performance criteria (2 m s^{-1}) proposed by Emery et al. (2001). The large positive bias in wind speed was also reported in other studies (Zhang et al., 2010; Gao et al., 2015; Liao et al., 2015; Qiu et al., 2017). The predicted J[NO_2] agrees well with the observations with R of 0.97 and

NMB of 6.8%. We also conduct comparison between observed and simulated temperature profiles at 08:00 and 20:00 LST in Beijing during 29 July to 1 August 2014 in Figure S2~~1~~. The vertical profile of observed temperature, especially the thermal inversion layer occurred on 31 July around 1600 m, is well captured by the model. Generally, the WRF-Chem model reasonably reproduces the temporal variations of observed meteorological parameters.

4 Results

It is known that co-occurrence of $PM_{2.5}$ and O_3 pollution is frequently observed nowadays over China (Dai et al., 2021). The complex air pollution characterized by high $PM_{2.5}$ and O_3 levels has already received widespread ~~attentions~~attention from both scientists and policy-makers. Therefore, we examine the impacts of aerosol-radiation interactions on O_3 concentrations with a special focus on the complex air pollution areas (CAPAs, Fig. ~~S2~~S3), where the mean simulated daily $PM_{2.5}$ and MDA8 (maximum daily 8-h average) O_3 concentrations are larger than 75 μg m^{-3} and 80 ppb, respectively, based on the National Ambient Air Quality Standards (http://www.mee.gov.cn).

4.1 Impacts of aerosol-radiation interactions on meteorology

Figure 4 shows the impacts of aerosol-radiation interactions on ~~downward~~ shortwave radiation at the surface (BOT_SW), ~~downward~~ shortwave radiation in the atmosphere (ATM_SW), PBLH, T_2, RH_2, and WS_{10} during the daytime (08:00-17:00 LST) from 28 July to 3 August 2014. As a result of the interactions between aerosol and radiation (the combined impacts of API and ARF), BOT_SW is decreased over the entire simulated domain. Over CAPAs, the BOT_SW is decreased by ~~115.8~~93.2 W m^{-2} (20.5%). Contrary to the changes in BOT_SW, ATM_SW is increased significantly with an increase of 72.8 W m^{-2} (25.3%) over CAPAs. The decreased BOT_SW perturbs the near-surface energy flux, which weakens convection and suppresses the development of PBL (Li et al., 2017b). The PBLH averaged over CAPAs is calculated to decrease by 129.0 m (13.0%). The reduced surface radiation budget can directly lead to changes in near-surface temperature. Therefore, the changes in T_2 have the similar

spatial patterns with BOT_SW; the surface temperature is decreased by 0.56 °C averaged over CAPAs. RH_2 is increased over most of the domain with an average rise of 2.4%, which is beneficial for the hygroscopic growth of aerosols. WS_{10} exhibits overall reductions over CAPAs and is calculated to decrease by 0.12 m s^{-1} on average. We also examine the changed meteorological variables caused by API and ARF respectively. As shown in Fig. ~~S3~~S4, API has little impact on meteorological variables; the above changes are mainly caused by ARF.

4.2 Impacts of aerosol-radiation interactions on photolysis

Figure 5 shows the spatial distribution of mean daytime surface $PM_{2.5}$ concentrations simulated by BASE case and the changes in J[NO_2] and J[O^1D] due to aerosol-radiation interactions from 28 July to 3 August 2014. When the combined impacts (API and ARF) are considered, J[NO_2] and J[O^1D] are decreased over the entire domain; the spatial patterns of changed J[NO_2] and J[O^1D] are similar to that of simulated $PM_{2.5}$. The surface J[NO_2] and J[O^1D] are decreased by 1.8×10^{-3} s^{-1} (40.5%) and 6.1×10^{-6} s^{-1} (48.8%) averaged over CAPAs. Figure ~~S4~~ S5 exhibits the percentage changes in surface J[NO_2] and J[O^1D] caused by API and ARF respectively. It is found that J[NO_2] and J[O^1D] are significantly modified by API and little affected by ARF.

4.3 Impacts of aerosol-radiation interactions on O_3

Figure 6 shows the changes in surface-layer O_3 due to API, ARF, and the combined effects (denoted as ALL). As shown in Fig. 6a, API alone leads to overall surface O_3 decreases over the entire domain with an average reduction of 8.5 ppb (10.1%) over CAPAs. The change can be explained by the substantially diminished UV radiation due to aerosol loading, which significantly weakens the efficiency of photochemical reactions and restrains O_3 formation. The decreased surface O_3 concentration due to ARF, however, is only 2.9 ppb (3.1%, Fig. 6b), which indicates that API is the dominant way for O_3 reduction related to aerosol-radiation interactions. ~~The distributions of changed O_3 concentrations coincide with NO_x variations (Fig. S5b). Since North China is VOC-limited (Jin et al., 2015), the increase in NO_x due to ARF may partly explain the O_3 decrease.~~ The combined effects of API and ARF are shown in Fig. 6c. Generally,

aerosol-radiation interactions decrease the surface O_3 concentration by 11.4 ppb (13.5%) averaged over CAPAs.

~~We further define an index to characterize the effects of aerosols on surface O_3 concentrations. The ratio of changes in O_3 to local $PM_{2.5}$ levels is defined as:~~

~~$$ROP = \frac{\Delta O_3}{PM_{2.5}\_BASE},$$~~

~~where ΔO_3 is the changed O_3 concentration caused by ALL, and $PM_{2.5}$_BASE is the surface $PM_{2.5}$ concentration simulated in the BASE scenario. The calculated ROP is -0.14 ppb ($\mu g\ m^{-3}$)$^{-1}$ averaged over CAPAs, which means when the concentrations of $PM_{2.5}$ is 100 $\mu g\ m^{-3}$, the O_3 decrease will be up to 14 ppb over CAPAs due to aerosol-radiation interactions.~~

4.4 Influencing mechanism of aerosol-radiation interactions on O_3

Figure 7a shows diurnal variations of simulated surface (first layer) daytime O_3 concentrations over CAPAs in three cases (BASE, NOAPI, and NOALL). All cases present O_3 increases from 08:00 LST. It is shown that the simulated O_3 concentrations in BASE case increase more slowly than that in NOAPI and NOALL cases. To explain the underlying mechanisms of API and ARF impacts on O_3, we quantify the variations in contributions of different processes (ADV, CHEM, and VMIX) to O_3 by using the IPR analysis.

Figure 7b shows hourly surface O_3 changes induced by each physical/chemical process (i.e., ADV, CHEM, and VMIX) in BASE case. The significant positive contribution to the hourly variation in O_3 is contributed by VMIX, and the contribution reaches the maximum at about 10:00 LST. After 14:00 LST, the contribution from VMIX remains constant (nearly +2 ppb h^{-1}), which is probably attributed to the stable boundary layer development (Tang et al., 2016). The CHEM process makes negative contributions at around 09:00 and 16:00 LST, which means that the chemical consumption of O_3 is stronger than the chemical production. At noon, the net chemical contribution turns to be positive due to stronger solar UV radiation. The contribution from all the processes (NET, the sum of VMIX, CHEM, and ADV) to O_3 is peaked at

the noon and then becomes weakened. After sunset (17:00 LST), the NET contribution turns to be negative over CAPAs, leading to O_3 decrease.

Figure 7c shows the changes in hourly process contributions caused by API. The chemical production of O_3 is suppressed significantly due to aerosol impacts on photolysis rates. The weakened O_3 chemical production decreases the contribution from CHEM, and results in a negative value of CHEM_DIF (-3.5 ppb h^{-1}). In contrast to CHEM_DIF, the contribution from changed VMIX (VMIX_DIF) to O_3 concentration due to API is always positive, and the mean value is +3.1 ppb h^{-1}. The impact of API on ADV process is relatively small (-0.36 ppb h^{-1}). NET_DIF, namely the sum of VMIX_DIF, CHEM_DIF and ADV_DIF, indicates the differences in hourly O_3 changes caused by API. As shown in Fig. 7c, NET_DIF is almost negative during the daytime over CAPAs with the mean value of -0.76 ppb h^{-1}. This is because the decreases in CHEM and ADV are larger than the increases in VMIX caused by API; the O_3 decrease is mainly attributed to the significantly decreased contribution from CHEM. The maximum difference in O_3 between BASE and NOAPI appears at 17:00 LST with a value of -10.1 ppb (Fig. 7a).

Figure 7d shows the impacts of ARF on each physical/chemical process contribution to the hourly O_3 variation. At 08:00 LST, the change in VMIX due to ARF is large with a value of -4.6 ppb h^{-1}, resulting in a net negative variation with all processes considered. The decrease in O_3 reaches the maximum with the value of 6.1 ppb at around 09:00 LST over CAPAs (Fig. 7a). During 10:00 to 16:00 LST, the positive VMIX_DIF (mean value of +0.59 ppb h^{-1}) or the positive CHEM_DIF (mean value of +0.16 ppb h^{-1}) is the major process to positive NET_DIF.

When both impacts of API and ARF are considered, the variation pattern of the difference in hourly process contribution shown in Fig. 7e is similar to that in Fig. 7c, which indicates that API is the dominant factor to surface-layer O_3 reduction.

Figure 8 presents the vertical profiles of simulated daytime O_3 concentrations in three cases (BASE, NOAPI, and NOALL), and the differences in contributions from each physical/chemical process to hourly O_3 variations caused by API, ARF and the

combined effects during 28 July to 3 August 2014 over CAPAs. As shown in Fig. 8a, the O_3 concentration is lower in BASE than that in other two scenarios (NOAPI and NOALL), especially at the lower 12 levels (below 731.9 m), owing to the impacts of aerosols (API and/or ARF).

The changes in each process contribution caused by API are presented in Fig. 8b. The contribution from CHEM_DIF is -2.14 ppb h^{-1} for first seven layers (from 23.4 to 290.7 m). Conversely, the contribution from VMIX_DIF shows a positive value under the 290.7 m (between first layer to seventh layer)~~at the lower seven layers~~ with the mean value of +1.7 ppb h^{-1}. The positive variation in VMIX due to API may be associated with the different vertical gradient of O_3 between BASE and NOAPI cases. The contributions of changed advections (ADVH_DIF and ADVZ_DIF) are relatively small, with mean values of +0.25 and -0.47 ppb h^{-1} respectively below the first seven layers, which may result from small impact of API on wind filed (Fig. ~~S3a~~S4a). The net difference is a negative value (-0.66 ppb h^{-1}); API leads to O_3 reduction not only nearly surface but also ~~in the~~ aloft.

Figure 8c shows the differences in O_3 budget due to ARF. When the ARF is considered, the vertical turbulence is weakened and the development of PBL is inhibited, which makes VMIX_DIF negative at the lower 7 layers (below the 290.7 m) with a mean value of -0.55 ppb h^{-1}, but the variation in CHEM caused by ARF is positive with a mean value of +0.6 ppb h^{-1}. ~~The chemical production of tropospheric O_3 is affected by both photolysis rate and the concentrations of precursors (Tie et al., 2009).~~ The enhanced O_3 precursors due to ARF can promote the chemical production of O_3 (Tie et al., 2009). The changes of ADVZ and ADVH (ADVZ_DIF and ADVH_DIF) caused by ARF are associated with the variations in wind filed. When ARF is considered, the horizontal wind speed is decreased (Fig. S6a), which makes ADVH_DIF positive at the lower twelve layers with a mean value of +0.5 ppb h^{-1}. However, ADVZ_DIF is negative at these layers with a mean value of -0.48 ppb h^{-1} because aerosol radiative effects decrease the transport of O_3 from the upper to lower layers (Fig. S6b).

In Fig. 8d, the pattern and magnitude of the differences in process contributions between

BASE and NOALL are similar to those caused by API, indicating again the dominate role of API on O_3 changes. The impacts of API on O_3 both near the surface and aloft are greater than those of ARF.

5 Discussions

In order to make the analysis and conclusions more robust, another two complex air pollution episodes (8-13 July 2015 and 5-11 June 2016) in this region are also selected to conduct simulations for generating general conclusions. Simulated air pollutants ($PM_{2.5}$ and O_3) and meteorological variables (T_2, RH_2, and WS_{10}) during 8-13 July 2015 (Episode 2) and 5-11 June 2016 (Episode 3) are compared with observations (Fig. S7-Fig. S8). In general, both the observed meteorological parameters and pollutant concentrations can be reasonably reproduced by the model, with correlation coefficients (R) of 0.56~0.98 and normalized mean bias (NMB) of –7.1%~+33.4%. More details about the model evaluation are listed in the supporting information (Text S1).

As shown in Fig. S9(a1-a2), API alone leads to the decrease in surface O_3 over the entire domain with an average reduction of 9.0 ppb (10.6%) and 8.3 ppb (10.4%) over CAPAs in Episode 2 and Episode 3, respectively. The decreased surface O_3 concentrations over CAPAs due to ARF are only 1.0 ppb (1.2%, Fig. 9(b1)) and 1.0 ppb (1.1%, Fig. 9(b2)) during Episode 2 and Episode 3, respectively. All the results indicate that API is the dominant factor for O_3 reduction related to aerosol-radiation interactions, the same as the conclusion analyzed from the case during 28 July to 3 August 2014. The combined effects of API and ARF decrease surface O_3 by 10.0 ppb (11.9%) and 9.3 ppb (11.6%) over CAPAs in Episode 2 and Episode 3, respectively. Analyzing Fig. S10 and Fig. S11, similar variation characteristics are shown in Episode 2 and Episode 3 as that during 28 July to 3 August 2014, with the larger impacts of API on O_3 both near the surface and aloft than those of ARF, indicating the dominant role of API on O_3 reduction related with aerosol-radiation interactions.

~~5~~ 6 Conclusions

In this study, the fully coupled regional chemistry transport model WRF-Chem is applied to investigate the impacts of aerosol-radiation interactions, including the impact of aerosol-photolysis interaction (API) and the impact of aerosol-radiation feedback (ARF), on O_3 during a summertime complex air pollution episode from 28 July to 3 August 2014. Three sensitivity experiments are designed to quantify the respective and combined impacts from API and ARF. Generally, the spatiotemporal distributions of observed pollutant concentrations and meteorological parameters are captured fairly well by the model with high correlation coefficients of 0.66–0.86 for pollutant concentrations and 0.70–0.98 for meteorological parameters.

Sensitivity experiments show that aerosol-radiation interactions decrease BOT_SW, T_2, WS_{10}, PBLH, J[NO_2], and J[O^1D] by ~~115.8~~93.2 W m^{-2}, 0.56 °C, 0.12 m s^{-1}, 129 m, 1.8×10^{-3} s^{-1}, and 6.1×10^{-6} s^{-1} over CAPAs, and increase ATM_SW and RH_2 by 72.8 W m^{-2} and 2.4%. The changed meteorological variables and weakened photochemistry reaction further reduce surface-layer O_3 concentration by up to 11.4 ppb (13.5%), with API and ARF contributing 74.6% and 25.4%, respectively. ~~The combined impacts of API and ARF on O_3 can be characterized by the ratio of changed O_3 (ΔO_3) to local $PM_{2.5}$ level ($PM_{2.5}$_BASE), defining as ROP = $\Delta O_3/PM_{2.5}$_BASE. The calculated ROP is ~0.14 ppb (μg m^{-3})$^{-1}$ averaged over CAPAs.~~

We further examine the influencing mechanism of aerosol-radiation interactions on O_3 by using integrated process rate analysis. API can directly affect O_3 by reducing the photochemistry reactions within the lower several hundred meters and therefore amplify the O_3 vertical gradient, which promotes ~~the contribution from VMIX and~~ the vertical mixing of O_3. The reduced photochemistry reactions of O_3 weaken the chemical contribution and reduce surface O_3 concentrations, even though the enhanced vertical mixing can partly counteract the reduction. ARF affects O_3 concentrations indirectly through the changed meteorological variables, e.g., the decreased PBLH. The suppressed PBL can weaken the vertical mixing of O_3 by turbulence. Generally, the impacts of API on O_3 both near the surface and aloft are greater than those of ARF, indicating the dominant role of API on O_3 reduction related with aerosol-radiation

interactions.

This study provides a detailed understanding of aerosol impacts on O_3 through aerosol-radiation interactions (including both API and ARF). The results imply that future $PM_{2.5}$ reductions will lead to O_3 increases due to weakened aerosol-radiation interactions. A recent study~~Recent study~~ emphasized the need for controlling VOCs emissions to mitigate O_3 pollution (Li et al., 2019). Therefore, tighter controls of O_3 precursors (especially VOCs emissions) are needed to counteract future O_3 increases caused by weakened aerosol-radiation interactions~~.~~, and the contributions of different mitigation strategies with the impacts of aerosol-radiation interactions to O_3 air quality will be discussed detailedly in our future work.

There are some limitations to this work. The uncertainty of the lack of secondary organic aerosols (SOA), and the missing mechanisms of some heterogeneous reactions may result in large uncertainties in the final simulation results. Gao et al. (2017) added some SOA formation mechanisms into the MOSAIC module by using the volatility basis set (VBS) in WRF-Chem and found that the surface $PM_{2.5}$ concentrations in urban Beijing were reduced by 1.9 $\mu g\ m^{-3}$ due to the weakened ARF effect during Asia-Pacific Economic Cooperation (APEC). Similar magnitude can also be found in Zhou et al. (2019) (-1.8 $\mu g\ m^{-3}$) who did not consider the impacts of SOA in WRF-Chem when analyzing the impacts of weakened ARF on $PM_{2.5}$ during APEC. Therefore, more work should be conducted to explore the impacts of ARF on $PM_{2.5}$ and O_3 concentrations under consideration of SOA in future.

Data availability

The observed hourly surface concentrations of air pollutants are derived from the China National Environmental Monitoring Center (http://www.cnemc.cn). The observed surface meteorological data are obtained from NOAA's National Climatic Data Center (https://gis.ncdc.noaa.gov/maps/ncei/cdo/hourly). The radiosonde data are provided by the University of Wyoming (http://weather.uwyo.edu/). The photolysis rates of nitrogen dioxide in Beijing are provided by Xin Li (li_xin@pku.edu.cn). The MODIS data are obtained from the NASA Level 1 and Atmosphere Archive and Distribution System (https://ladsweb.modaps.eosdis.nasa.gov). The simulation results can be accessed by contacting Lei Chen (chenlei@nuist.edu.cn) and Hong Liao (hongliao@nuist.edu.cn).

Author contributions

HY, LC, and HL conceived the study and designed the experiments. HY and LC performed the simulations and carried out the data analysis. JZ, WW, and XL provided useful comments on the paper. HY prepared the paper with contributions from all co-authors.

Competing interests

The authors declare that they have no competing interests.

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Reference

Albrecht, B. A.: Aerosols, cloud microphysics, and fractional cloudiness, Science, 245, 1227–1230, 1989.

Chen, F. and Dudhia, J.: Coupling an Advanced Land Surface – Hydrology Model with the Penn State – NCAR MM5 Modeling System. Part I: Model Implementation and Sensitivity, Mon. Weather Rev., 129(4), 569–585, 2001.

Chen, L., Zhu, J., Liao, H., Gao, Y., Qiu, Y., Zhang, M., Liu, Z., Li, N., and Wang, Y.: Assessing the formation and evolution mechanisms of severe haze pollution in the Beijing–Tianjin–Hebei region using process analysis, Atmos. Chem. Phys., 19, 10845–10864, https://doi.org/10.5194/acp-19-10845-2019, 2019.

Cheng, Y., Zheng, G., Chao, W., Mu, Q., Bo, Z., Wang, Z., Meng, G., Qiang, Z., He, K., and Carmichael, G.: Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events in China, Science Advances, 2, https://doi.org/10.1126/sciadv.1601530, 2016.

Dai, H., Zhu, J., Liao, H., Li, J., Liang, M., Yang, Y., and Yue, X.: Co-occurrence of ozone and $PM_{2.5}$ pollution in the Yangtze River Delta over 2013–2019: Spatiotemporal distribution and meteorological conditions, Atmos. Res., 249, 105363, 2021.

Dickerson, R. R., Kondragunta, S., Stenchikov, G., Civerolo, K. L., Doddridge, B. G., and Holben, B. N.: The impact of aerosols on solar ultraviolet radiation and photochemical smog, Science, 278, 827-830, 10.1126/science.278.5339.827, 1997.

Emery, C., Tai, E., and Yarwood, G.: Enhanced meteorological modeling and performance evaluation for two Texas ozone episodes, in: Prepared for the Texas Natural Resource Conservation Commission, ENVIRON International Corporation, Novato, CA, USA, 2001.

Emmons, L. K., Walters, S., Hess, P. G., Lamarque, J.-F., Pfister, G. G., Fillmore, D., Granier, C., Guenther, A., Kinnison, D., Laepple, T., Orlando, J., Tie, X., Tyndall, G., Wiedinmyer, C., Baughcum, S. L., and Kloster, S.: Description and evaluation of the Model for Ozone and Related chemical Tracers, version 4 (MOZART-4), Geosci.

Model Dev., 3, 43–67, doi:10.5194/gmd- 3-43-2010, 2010.

Foken, T.: 50 years of the Monin-Obukhov similarity theory, Bound.-Layer Meteor., 119, 431–437, 2006.

Gao, J., Li, Y., Zhu, B., Hu, B., Wang, L., and Bao, F.: What have we missed when studying the impact of aerosols on surface ozone via changing photolysis rates?, Atmos. Chem. Phys., 20, 10831–10844, https://doi.org/10.5194/acp-20-10831-2020, 2020.

Gao, J. H., Zhu, B., Xiao, H., Kang, H. Q., Pan, C., Wang, D. D., and Wang, H. L.: Effects of black carbon and boundary layer interaction on surface ozone in Nanjing, China, Atmos. Chem. Phys., 18, 7081–7094, https://doi.org/10.5194/acp-18-7081-2018, 2018.

Gao, M., Carmichael, G. R., Wang, Y., Saide, P. E., Yu, M., Xin, J., Liu, Z., and Wang, Z.: Modeling study of the 2010 regional haze event in the North China Plain, Atmos. Chem. Phys., 16, 1673–1691, doi:10.5194/acp-16-1673-2016, 2016a.

Gao, J., Zhu, B., Xiao, H., Kang, H., Hou, X., and Shao, P.: A case study of surface ozone source apportionment during a high concentration episode, under frequent shifting wind conditions over the Yangtze River Delta, China, Sci. Total Environ., 544, 853, https://doi.org/10.1016/j.scitotenv.2015.12.039, 2016b.

Gao, M., Liu, Z., Wang, Y., Lu, X., Ji, D. and Wang, L.: Distinguishing the roles of meteorology, emission control measures, regional transport, and co-benefits of reduced aerosol feedbacks in “APEC” Blue, Atmos. Environ., 167, 476–486, 420 doi:10.1016/j.atmosenv.2017.08.054, 2017.

Gao, Y., Zhang, M., Liu, Z., Wang, L., Wang, P., Xia, X., Tao, M., and Zhu, L.: Modeling the feedback between aerosol and meteorological variables in the atmospheric boundary layer during a severe fog–haze event over the North China Plain, Atmos. Chem. Phys., 15, 4279–4295, doi:10.5194/acp-15-4279-2015, 2015.

Goncalves, M., Jimenez-Guerrero, P., Baldasano, J.M.: Contribution of atmospheric processes affecting the dynamics of air pollution in South-Western Europe during a typical summertime photochemical episode, Atmos. Chem. Phys., 9, 849–864,

doi:10.5194/acp-9-849-2009, 2009.

Grell, G. A., Peckham, S. E., Schmitz, R., Mckeen, S. A., Frost, G., Skamarock, K., and Eder, B.: Fully coupled "online" chemistry within the WRF model, Atmos. Environ., 39, 6957–6975, 2005.

Guenther, A., Karl, T., Harley, P., Wiedinmyer, C., Palmer, P. I., and Geron, C.: Estimates of global terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols from Nature), Atmos. Chem. Phys., 6, 3181–3210, doi:10.5194/acp-6-3181-2006, 2006.

Hansen, J., Sato, M., and Ruedy, R.: Radiative forcing and climate response, J. Geophys. Res., 102, 6831–6864, 1997.

Haywood, J. and Boucher, O.: Estimates of the direct and indirect radiative forcing due to tropospheric aerosols: A review, Rev. Geophys., 38, 513–543, 2000.

Hong, S.-Y., Noh, Y., and Dudhia, J.: A New Vertical Diffusion Package with an Explicit Treatment of Entrainment Processes, Mon. Weather Rev., 134, 2318–2341, 2006.

Iacono, M. J., Delamere, J. S., Mlawer, E. J., Shephard, M. W., Clough, S. A., and Collins, W. D.: Radiative forcing by long-lived greenhouse gases: Calculations with the AER radiative transfer models, J. Geophys. Res., 113, D13103, doi:10.1029/2008JD009944, 2008.

~~Jin, X. and Holloway, T.: Spatial and temporal variability of ozone sensitivity over China observed from the Ozone Monitoring Instrument, J. Geophys. Res. Atmos., 120, 7229–7246, https://doi.org/10.1002/2015JD023250, 2015.~~

Li, G., Bei, N., Tie, X., and Molina, L. T.: Aerosol effects on the photochemistry in Mexico City during MCMA-2006/MILAGRO campaign, Atmos Chem Phys, 11, 5169-5182, 10.5194/acp-11-5169-2011, 2011.

Li, K., Jacob, D. J., Liao, H., Zhu, J., Shah, V., Shen, L., Bates, K. H., Zhang, Q., and Zhai, S.: A two-pollutant strategy for improving ozone and particulate air quality in China, Nat. Geosci., 12, 906–910, https://doi.org/10.1038/s41561-019-0464-x, 2019.

Li, M., Zhang, Q., Kurokawa, J.-I., Woo, J.-H., He, K., Lu, Z., Ohara, T., Song, Y.,

559 Streets, D. G., Carmichael, G. R., Cheng, Y., Hong, C., Huo, H., Jiang, X., Kang, S.,
560 Liu, F., Su, H., and Zheng, B.: MIX: a mosaic Asian anthropogenic emission
561 inventory under the international collaboration framework of the MICS-Asia and
562 HTAP, Atmos. Chem. Phys., 17, 935–963, https://doi.org/10.5194/acp-17-935-2017,
563 2017a.
564 Li, Z., Guo, J., Ding, A., Liao, H., Liu, J., Sun, Y., Wang, T., Xue, H., Zhang, H., and
565 Zhu, B.: Aerosol and boundary-layer interactions and impact on air quality, Nat. Sci.
566 Rev., 4, 810–833, https://doi.org/10.1093/nsr/nwx117, 2017b.
567 Liao, H., Yung, Y. L., and Seinfeld, J. H.: Effects of aerosols on tropospheric photolysis
568 rates in clear and cloudy atmospheres, J. Geophys. Res., 104, 23697–23707, 1999.
569 Liao, L., Lou, S. J., Fu, Y., Chang, W. Y., and Liao, H.: Radiative forcing of aerosols
570 and its impact on surface air temperature on the synoptic scale in eastern China [in
571 Chinese], Chin. J. Atmos. Sci., 39, 68–82, doi: 10.3878/j.issn.1006-9895.1402.13302,
572 2015.
573 Lin, Y.-L., Farley, R. D., and Orville, H. D.: Bulk parameterization of the snow field in
574 a cloud model, J. Clim. Appl. Meteorol., 22, 1065–1092, 1983.
575 Liu, Y. and Wang, T.: Worsening urban ozone pollution in China from 2013 to 2017 –
576 Part 1: The complex and varying roles of meteorology, Atmos. Chem. Phys., 20,
577 6305–6321, https://doi.org/10.5194/acp-20-6305-2020, 2020.
578 Lo, J. C.-F., Yang, Z. L., and Pielke Sr, R. A.: Assessment of three dynamical climate
579 downscaling methods using the Weather Research and Forecasting (WRF) model, J.
580 Geophys. Res., 113, D09112, doi:10.1029/2007jd009216, 2008.
581 Lohmann, U., and Feichter, J.: Global indirect aerosol effects: A review. Atmospheric
582 Chemistry and Physics, 5, 715–737, https://doi.org/10.5194/acp-5-715-2005, 2005.
583 Lou, S., Liao, H., and Zhu, B.: Impacts of aerosols on surface-layer ozone
584 concentrations in China through heterogeneous reactions and changes in photolysis
585 rates, Atmos. Environ., 85, 123–138, 2014.
586 Miao, Y., Liu, S., Guo, J., Huang, S., Yan, Y., and Lou, M.: Unraveling the relationships
587 between boundary layer height and $PM_{2.5}$ pollution in China based on four-year

588 radiosonde measurements, *Environmental Pollution*,
589 https://doi.org/10.1016/j.envpol.2018.09.070, 2018.
590 Otte, T. L.: The impact of nudging in the meteorological model for retrospective air
591 quality simulations. Part I: Evaluation against national observation networks. J. Appl.
592 Meteor. Climatol., 47, 1853–1867, 2008.
593 Petäjä, T., Järvi, L., Kerminen, V. M., Ding, A. J., Sun, J. N., Nie, W., Kujansuu, J.,
594 Virkkula, A., Yang, X., Fu, C. B., Zilitinkevich, S., and Kulmala, M.: Enhanced air
595 pollution via aerosol-boundary layer feedback in China, Sci. Rep., 6, 18998,
596 doi:10.1038/srep18998, 2016.
597 Qiu, Y., Liao, H., Zhang, R., and Hu, J.: Simulated impacts of direct radiative effects
598 of scattering and absorbing aerosols on surface layer aerosol concentrations in China
599 during a heavily polluted event in February 2014, J. Geophys. Res. Atmos., 122,
600 5955–5975, doi:10.1002/2016JD026309, 2017.
601 Skamarock, W., Klemp, J. B., Dudhia, J., Gill, D. O., Barker, D. M., Duda, M., Huang,
602 X. Y., Wang,W., and Powers, J. G.: A description of the advanced research WRF
603 version 3, NCAR technical note NCAR/TN/u2013475, 2008.
604 Tang, G. Q., Zhu, X.W., Xin, J. Y., Hu, B., Song, T., Sun, Y., Zhang, J. Q., Wang, L. L.,
605 Cheng, M. T., Chao, N., Kong, L. B., Li, X., and Wang, Y. S.: Modelling study of
606 boundary-layer ozone over northern China – Part I: Ozone budget in summer, Atmos.
607 Res., 187, 128–137, 2017.
608 Tang, G., Zhang, J., Zhu, X., Song, T., Münkel, C., Hu, B., Schäfer, K., Liu, Z., Zhang,
609 J., Wang, L., Xin, J., Suppan, P., and Wang, Y.: Mixing layer height and its
610 implications for air pollution over Beijing, China, Atmos. Chem. Phys., 16, 2459–
611 2475, doi:10.5194/acp-16-2459-2016, 2016.
612 Tie, X., Geng, F., Li, P., Gao, W., and Zhao, C.: Measurement and modelling of ozone
613 variability in Shanghai, China, Atmos. Environ., 43, 4289–4302, 2009.
614 Wang, W., Li, X., Shao, M., Hu, M., Zeng, L., Wu, Y., and Tan, T.: The impact of
615 aerosols on photolysis frequencies and ozone production in Beijing during the 4-year
616 period 2012–2015, Atmos. Chem. Phys., 19, 9413–9429,

https://doi.org/10.5194/acp19-9413-2019, 2019.

Wild, O., Zhu, X., and Prather, M. J.: Fast-J: Accurate simulation of in- and below-cloud photolysis in tropospheric chemical models, J. Atmos. Chem., 37, 245–282, doi:10.1023/A:1006415919030, 2000.

Wu, J., Bei, N., Hu, B., Liu, S., Wang, Y., Shen, Z., Li, X., Liu, L., Wang, R., Liu, Z., Cao, J., Tie, X., Molina, L. T., Li, G.: Aerosol-photolysis interaction reduces particulate matter during wintertime haze events, Proc. Natl. Acad. Sci. USA, 117, 9755–9761, 2020.

Xing, J., Wang, J. D., Mathur, R., Wang, S. X., Sarwar, G., Pleim, J., Hogrefe, C., Zhang, Y. Q., Jiang, J. K., Wong, D. and Hao, J. M.: Impacts of aerosol direct effects on tropospheric ozone through changes in atmospheric dynamics and photolysis rates, Atmos. Chem. Phys., 17, 9869–9883, https://doi.org/10.5194/acp-17-9869-2017, 2017.

Zaveri, R. A. and Peters, L. K.: A new lumped structure photochemical mechanism for large-scale applications, J. Geophys. Res., 104, D23, 30387–30415, https://doi.org/10.1029/1999JD900876, 1999.

Zaveri, R. A., Easter, R. C., Fast, J. D., and Peters, L. K.: Model for simulating aerosol interactions and chemistry (MOSAIC), J. Geophys. Res., 113, D13204, https://doi.org/10.1029/2007JD008782, 2008.

Zhang, X., Zhang, Q., Hong, C. P., Zheng, Y. X., Geng, G. N., Tong, D., Zhang, Y. X., and Zhang, X. Y.: Enhancement of $PM_{2.5}$ concentrations by aerosol-meteorology interactions over China. Journal of Geophysical Research: Atmospheres,123, 1179–1194, https://doi.org/10.1002/2017JD027524, 2018.

Zhang, Y., Wen, X.-Y., and Jang, C. J.: Simulating chemistry– aerosol–cloud–radiation–climate feedbacks over the continental US using the online-coupled Weather Research Forecasting Model with chemistry (WRF/Chem), Atmos. Environ., 44, 3568–3582, doi:10.1016/j.atmosenv.2010.05.056, 2010.

Zhao, H.; Zheng, Y., and Li, C. Spatiotemporal distribution of $PM_{2.5}$ and O_3 and their interaction during the summer and winter seasons in Beijing, China. Sustainability,

10, 4519, 2018.

Zhou, M., Zhang, L., Chen, D., Gu, Y., Fu, T.-M., Gao, M., Zhao, Y., Lu, X. and Zhao, B.: The impact of aerosol-radiation interactions on the effectiveness of emission control measures, Environmental Research Letters, 14(2), 024002, https://doi.org/10.1088/1748-9326/aaf27d, 2019.

Zhu, J., Chen, L., Liao, H., and Dang, R.: Correlations between $PM_{2.5}$ and Ozone over China and Associated Underlying Reasons, Atmosphere, 352, 1–15, https://doi.org/10.3390/atmos10070352, 2019.

Zhu, J., Chen, L., Liao, H., Yang, H., Yang, Y., and Yue, X.: Enhanced $PM_{2.5}$ Decreases and O_3 Increases in China During COVID-19 Lockdown by Aerosol-Radiation Feedback, Geophys. Res. Lett., 48, https://doi.org/10.1029/2020GL090260, 2021.

Table 1. Physical parameterization options used in the simulation.

Options	Schemes
Microphysics scheme	Lin (Purdue) scheme (Lin et al.,1983)
Cumulus scheme	Grell 3D ensemble scheme
Boundary layer scheme	Yonsei University PBL scheme (Hong et al., 2006)
Surface layer scheme	Monin-Obukhov surface scheme (Foken, 2006)
Land-surface scheme	Unified Noah land-surface model (Chen and Dudhia, 2001)
Longwave radiation scheme	RRTMG (Iacono et al., 2008)
Shortwave radiation scheme	RRTMG (Iacono et al., 2008)

Table 2. Statistical parameters between simulated and observed $PM_{2.5}$ (μg m^{-3}), O_3 (ppb), 2 m temperature (T_2, °C), 2 m relative humidity (RH_2, %), 10 m wind speed (WS_{10}, m s^{-1}), and photolysis rate of NO_2 ($J[NO_2]$, s^{-1}) during 28 July to 3 August 2014.

Variables	O[a]	M[a]	R[b]	MB[c]	ME[d]	NMB[e](%)	NME[f](%)	RMSE[g]
$PM_{2.5}$	113.3	90.7	0.66	-21.8	25.2	-19.2	22.2	30.1
O_3	47.7	44.1	0.86	-5.7	15.5	-12.0	32.4	18.2
T_2	28.4	28.0	0.98	~~-0.2~~1.5	~~0.9~~1.6	~~-0.7~~5.7	~~3.3~~5.8	~~1.1~~1.8
RH_2	70.9	65.7	~~0.93~~ 0.91	~~-6.0~~0.5	~~6.7~~5.3	~~-8.5~~0.7	~~9.5~~7.9	~~8.7~~7.0
WS_{10}	2.4	3.0	~~0.70~~ 0.89	~~0.6~~0.7	~~0.9~~0.8	~~27.9~~28.5	~~36.6~~32.1	~~1.0~~0.9
$J[NO_2]$	1.6×10^{-3}	1.8×10^{-3}	0.97	1.1×10^{-4}	3×10^{-4}	6.8	18.5	5.3×10^{-4}

[a]***O*** and ***M*** are the averages for observed and simulated results, respectively. $O = \frac{1}{n}\times\sum_{i=1}^{n} O_i$, $M = \frac{1}{n}\times\sum_{i=1}^{n} M_i$.

[b]***R*** is the correlation coefficient between observations and model results. $R = \frac{\sum_{i=1}^{n}|(O_i-O)\times(M_i-M)|}{\sqrt{\sum_{i=1}^{n}(O_i-O)^2\times\sum_{i=1}^{n}(M_i-M)^2}}$.

[c]***MB*** is the mean bias between observations and model results. $MB = \frac{1}{n}\times\sum_{i=1}^{n}(M_i-O_i)$.

[d]***ME*** is the mean error between observations and model results. $ME = \frac{1}{n}\times\sum_{i=1}^{n}|M_i-O_i|$.

[e]***NMB*** is the normalized mean bias between observations and model results. $NMB = \frac{1}{n}\times\sum_{i=1}^{n}\frac{M_i-O_i}{O_i}\times100\%$.

[f]***NME*** is normal mean error between observations and model results. $NME = \frac{1}{n}\times\sum_{i=1}^{n}\frac{|M_i-O_i|}{O_i}\times100\%$.

[g]***RMSE*** is the root-mean-square error of observations and model results. $RMSE = \sqrt{\frac{1}{n}\times\sum_{i=1}^{n}(M_i-O_i)^2}$.

In the above O_i and M_i are the hourly observed and simulated data, respectively, and n is the total number of hours.

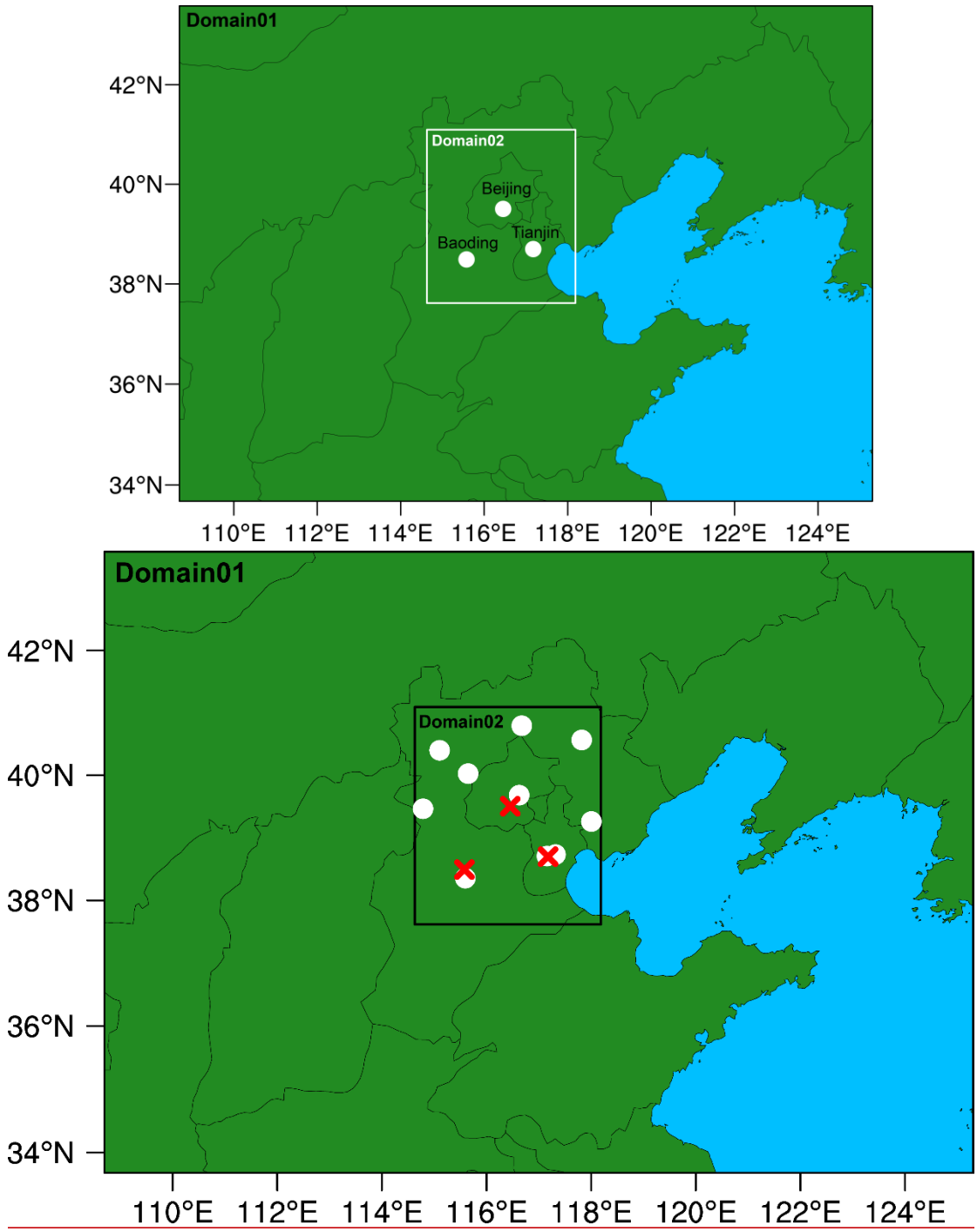
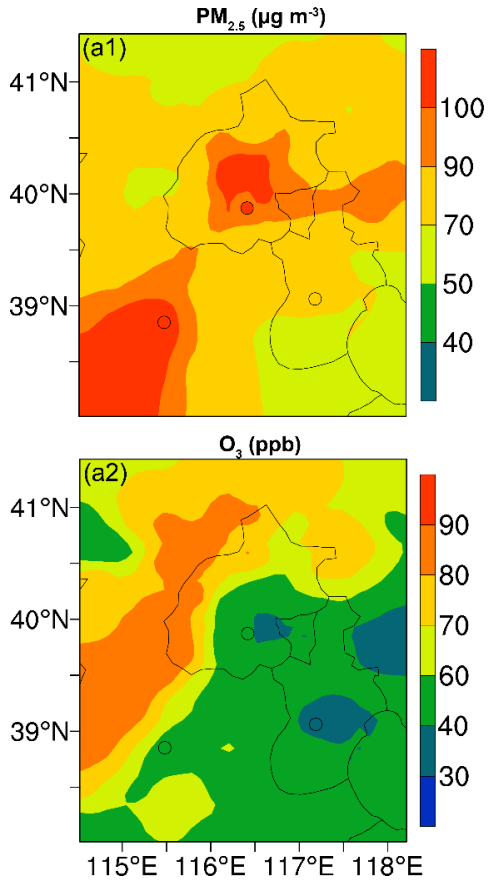
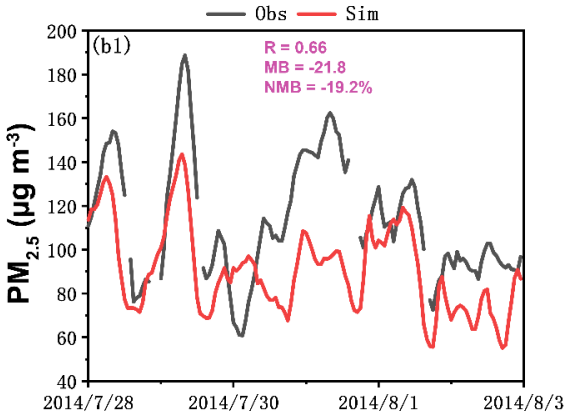
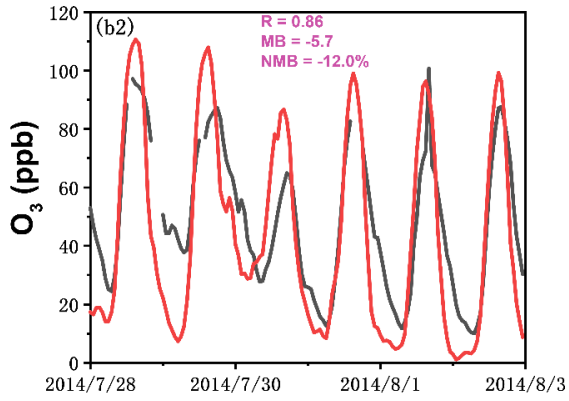


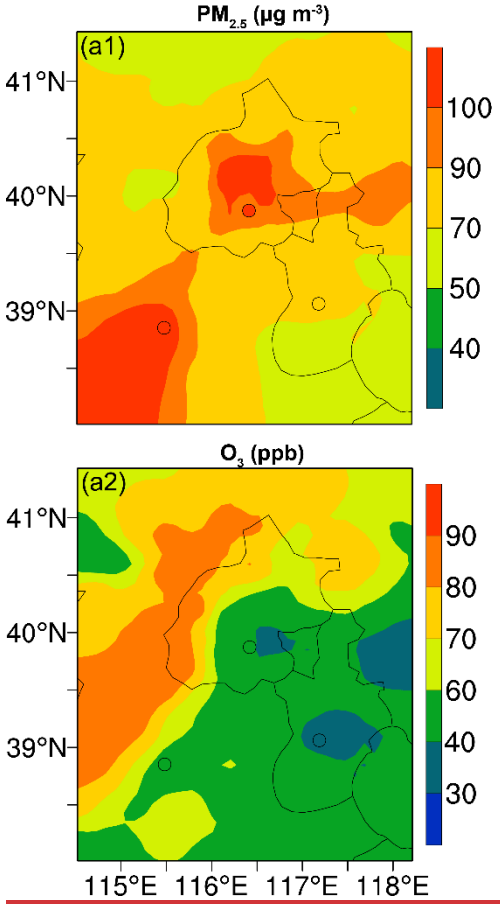
Figure 1. Map of the two WRF-Chem modeling domains with the locations of meteorological (white dots) and environmental (red crosses) observation sites used for model evaluation.~~Map of the two WRF-Chem modeling domains and the locations of observation sites (white dots) used for model evaluation.~~

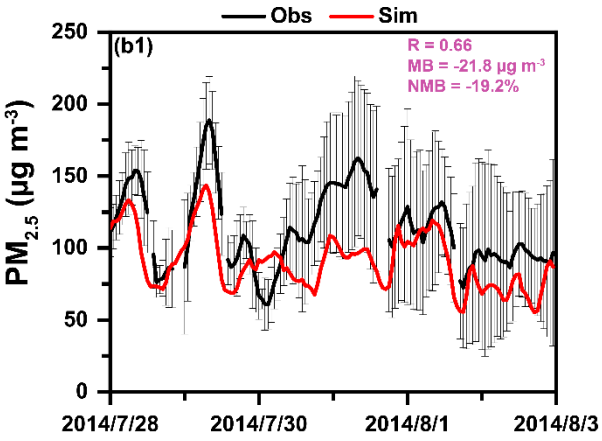

PM2.5 (μg m-3)
(a1)
41°N
40°N
39°N
100
90
70
50
40
O3 (ppb)
(a2)
90
80
70
60
40
30
115°E 116°E 117°E 118°E

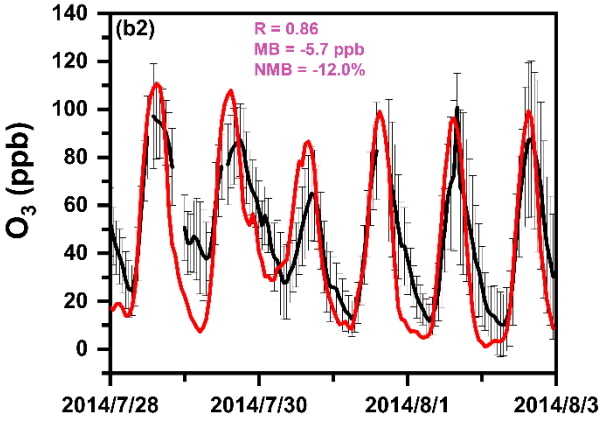

Obs
Sim
(b1)
R = 0.66
MB = -21.8
NMB = -19.2%
PM2.5 (μg m-3)
2014/7/28
2014/7/30
2014/8/1
2014/8/3


(b2)
R = 0.86
MB = -5.7
NMB = -12.0%
O3 (ppb)
2014/7/28
2014/7/30
2014/8/1
2014/8/3

1

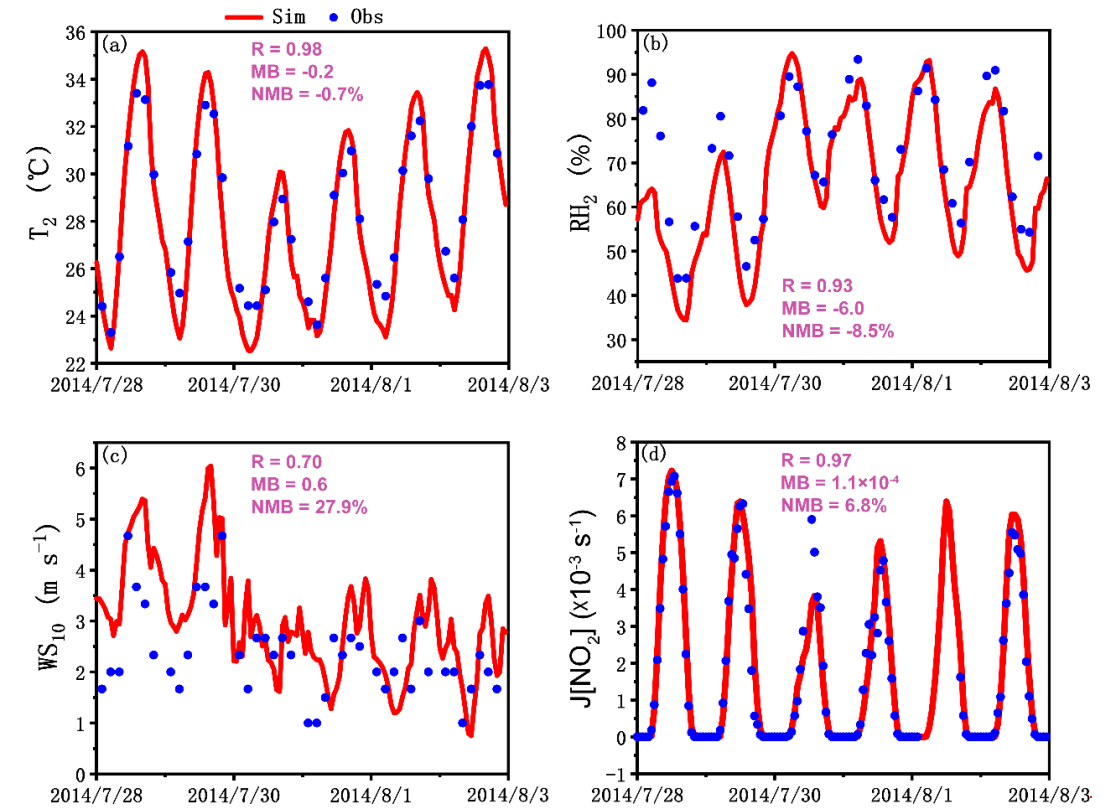

PM2.5 (μg m-3)
(a1)
41°N
40°N
39°N
100
90
70
50
40
O3 (ppb)
(a2)
90
80
70
60
40
30
115°E 116°E 117°E 118°E


Obs
Sim
(b1)
R = 0.66
MB = -21.8 μg m-3
NMB = -19.2%
PM2.5 (μg m-3)
2014/7/28
2014/7/30
2014/8/1
2014/8/3


(b2)
R = 0.86
MB = -5.7 ppb
NMB = -12.0%
O3 (ppb)
2014/7/28
2014/7/30
2014/8/1
2014/8/3

2

Figure 2. (a1-a2) Spatial distributions of simulated (color counters) and observed (colored circles) surface $PM_{2.5}$ and O_3 concentrations averaged during 28 July to 3 August 2014. (b1-b2) Time series of observed (black) and simulated (red) hourly surface $PM_{2.5}$ and O_3 concentrations averaged over the thirty-two~~32~~ observation sites in Beijing, Tianjin, and Baoding. The error bars in (b1) and (b2) are standard deviation on those average. The calculated correlation coefficient (R), mean bias (MB), and normalized mean bias (NMB) are also shown.



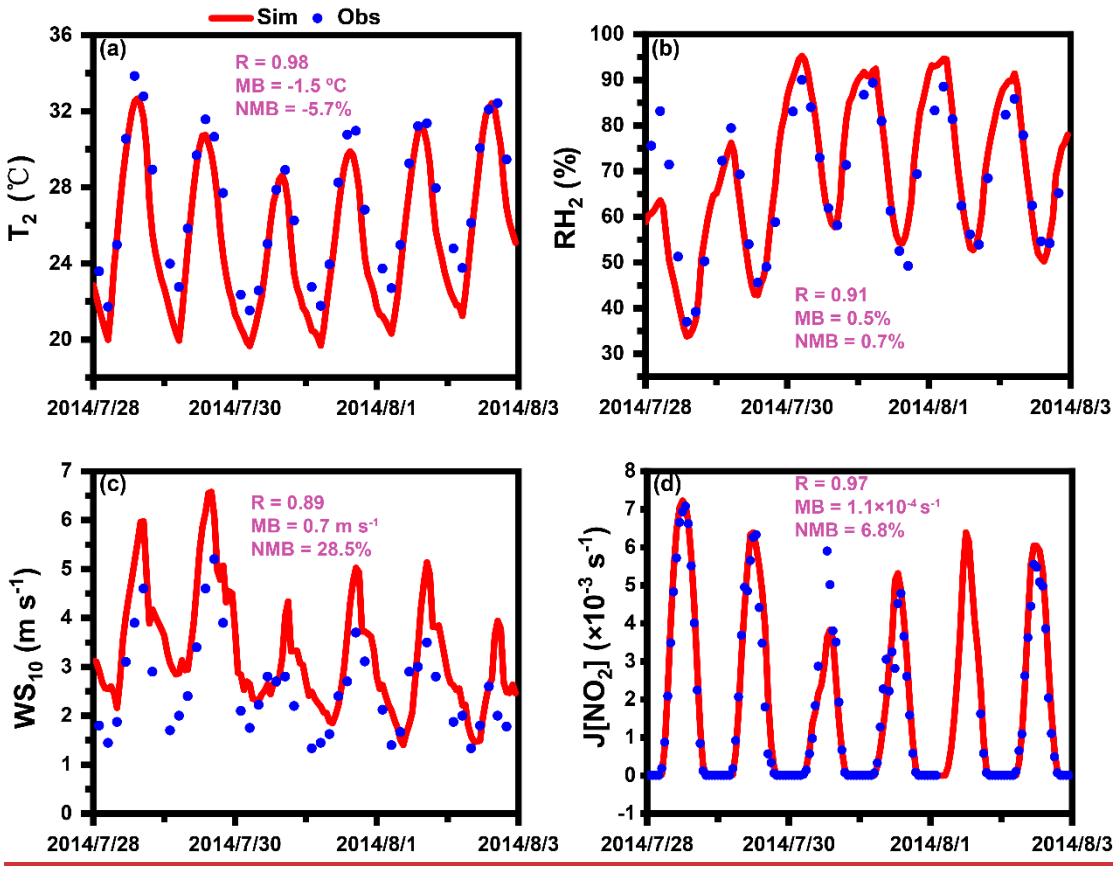
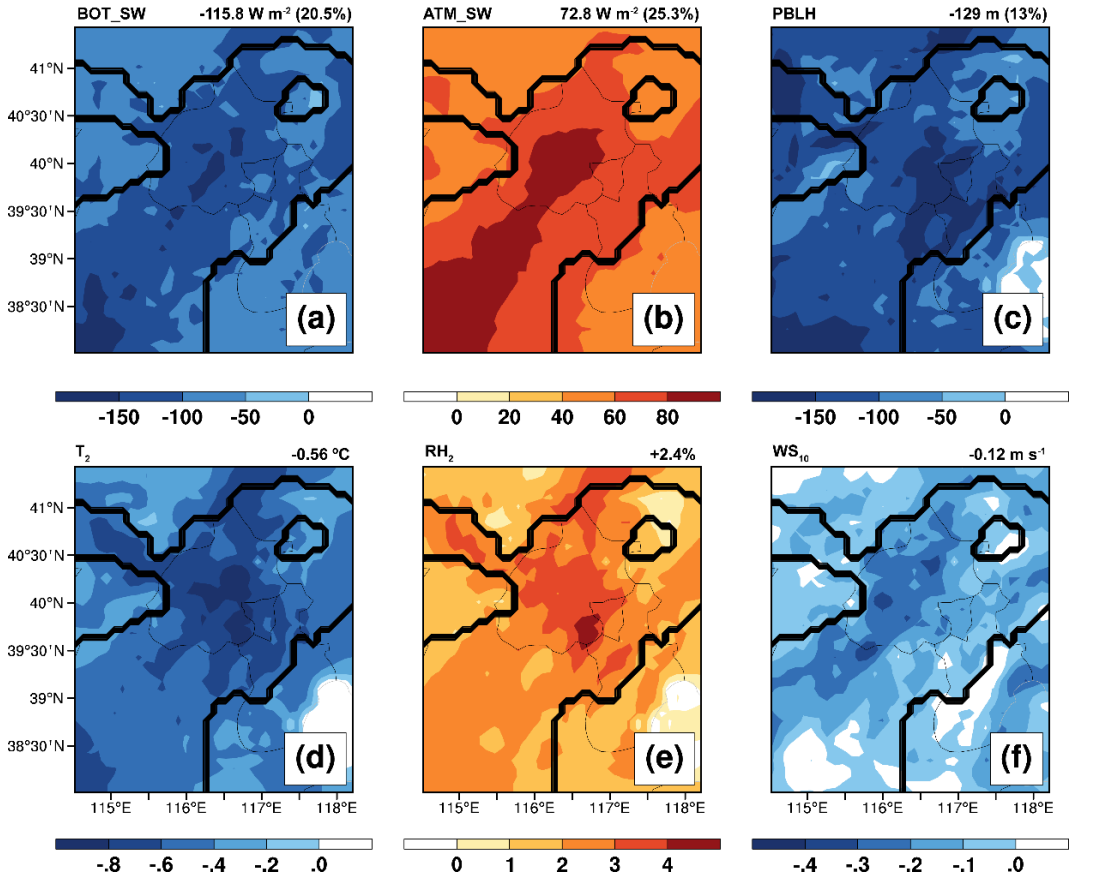
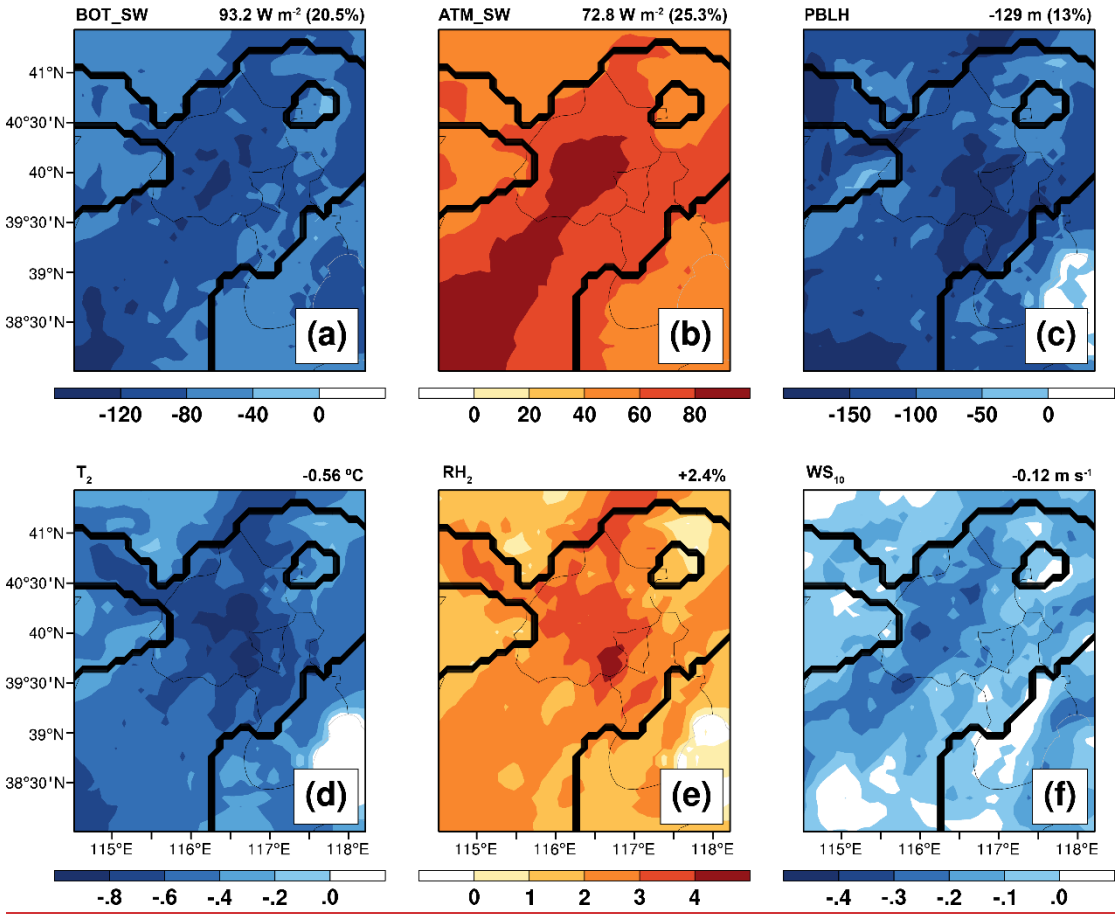


Figure 3. Time series of 3-hourly observed (blue dots) and hourly simulated (red lines) (a) 2-m temperature (T_2), (b) 2-m relative humidity (RH_2), (c) wind speed at 10 m (WS_{10}) averaged over ten meteorological observation stations, and (d) surface photolysis rate

of NO_2 ($J[NO_2]$) during 28 July to 3 August 2014. The calculated correlation coefficient (R), mean bias (MB), and normalized mean bias (NMB) are also shown.


BOT_SW -115.8 W m-2 (20.5%)
ATM_SW 72.8 W m-2 (25.3%)
PBLH -129 m (13%)
T2 -0.56 °C
RH2 +2.4%
WS10 -0.12 m s-1
41°N
40°30'N
40°N
39°30'N
39°N
38°30'N
115°E
116°E
117°E
118°E
(a)
(b)
(c)
(d)
(e)
(f)
-150 -100 -50 0
0 20 40 60 80
-150 -100 -50 0
-.8 -.6 -.4 -.2 .0
0 1 2 3 4
-.4 -.3 -.2 -.1 .0

1


BOT_SW 93.2 W m-2 (20.5%)
ATM_SW 72.8 W m-2 (25.3%)
PBLH -129 m (13%)
T2 -0.56 °C
RH2 +2.4%
WS10 -0.12 m s-1
41°N
40°30'N
40°N
39°30'N
39°N
38°30'N
115°E
116°E
117°E
118°E
(a)
(b)
(c)
(d)
(e)
(f)
-120 -80 -40 0
0 20 40 60 80
-150 -100 -50 0
-.8 -.6 -.4 -.2 .0
0 1 2 3 4
-.4 -.3 -.2 -.1 .0

2

Figure 4. The impacts of aerosol-radiation interactions on (a) ~~downward~~ shortwave radiation at the surface (BOT_SW), (b) ~~downward~~ shortwave radiation in the atmosphere (ATM_SW), (c) PBL height (PBLH), (d) 2-m temperature (T_2), (e) 2-m relative humidity (RH_2), and (f) 10-m wind speed (WS_{10}) during the daytime (08:00-17:00 LST) from 28 July to 3 August 2014. The region sandwiched between two black lines is defined as the complex air pollution areas (CAPAs) where the mean daily $PM_{2.5}$ and MDA8 O_3 concentrations in BASE case are larger than 75 μg m^{-3} and 80 ppb. The calculated changes averaged over CAPAs are also ~~shwon~~ shown at the top of each panel.

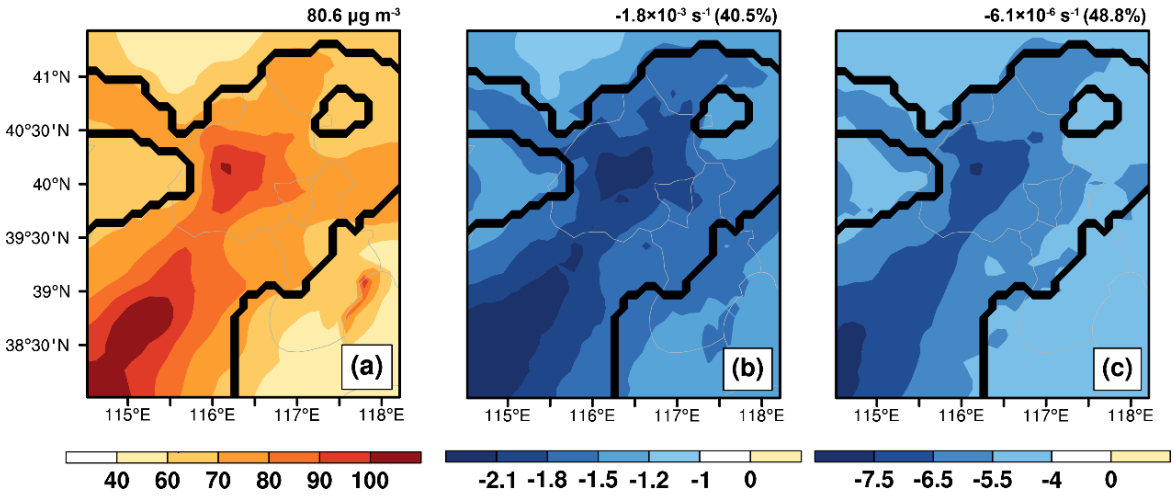


Figure 5. Spatial distributions of (a) simulated surface-layer $PM_{2.5}$ concentrations in BASE case, and changes in surface (b) J[NO_2] and (c) J[O^1D] due to aerosol-radiation interactions during the daytime (08:00-17:00 LST) from 28 July to 3 August 2014. The calculated values (percentage changes) avaraged over CAPAs are also ~~shwon~~ shown at the top of each panel.

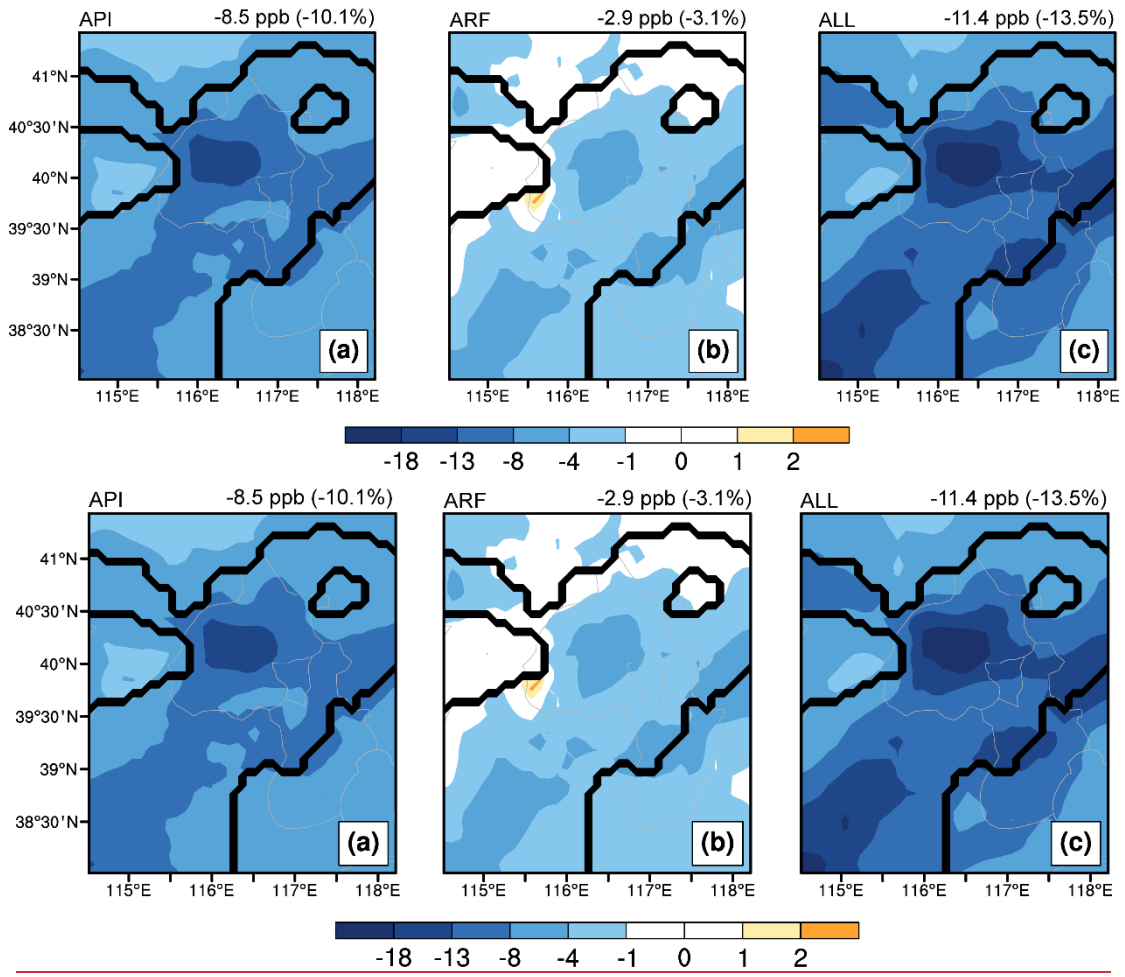


Figure 6. The changes in surface-layer ozone due to (a) aerosol-photolysis interaction (API), (b) aerosol-radiation feedback (ARF), and (c) the combined effects (ALL, defined as API+ARF) during the daytime (08:00-17:00 LST) from 28 July to 3 August 2014. The calculated mean changes avaraged over CAPAs are also shown at the top of each panel.

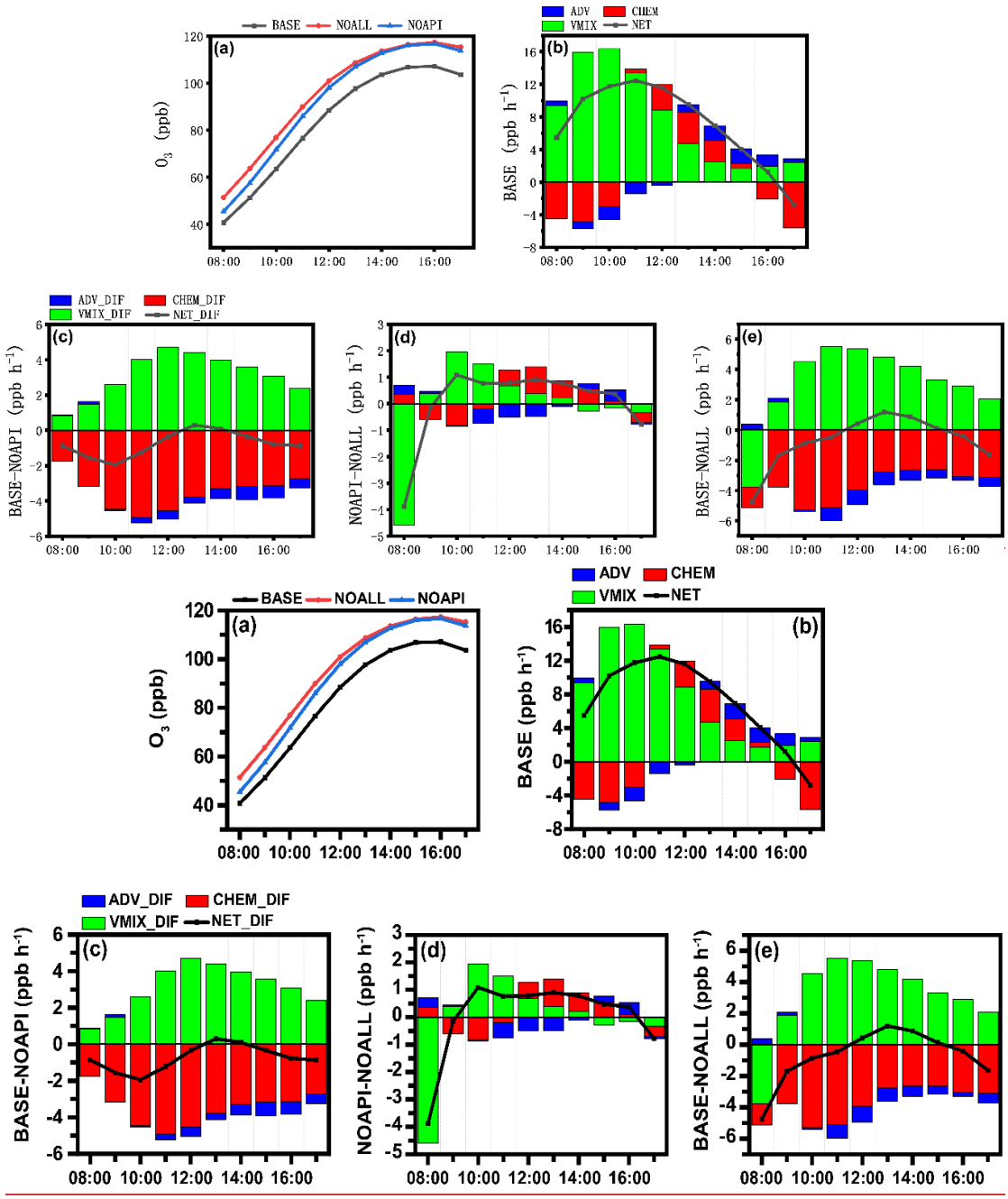
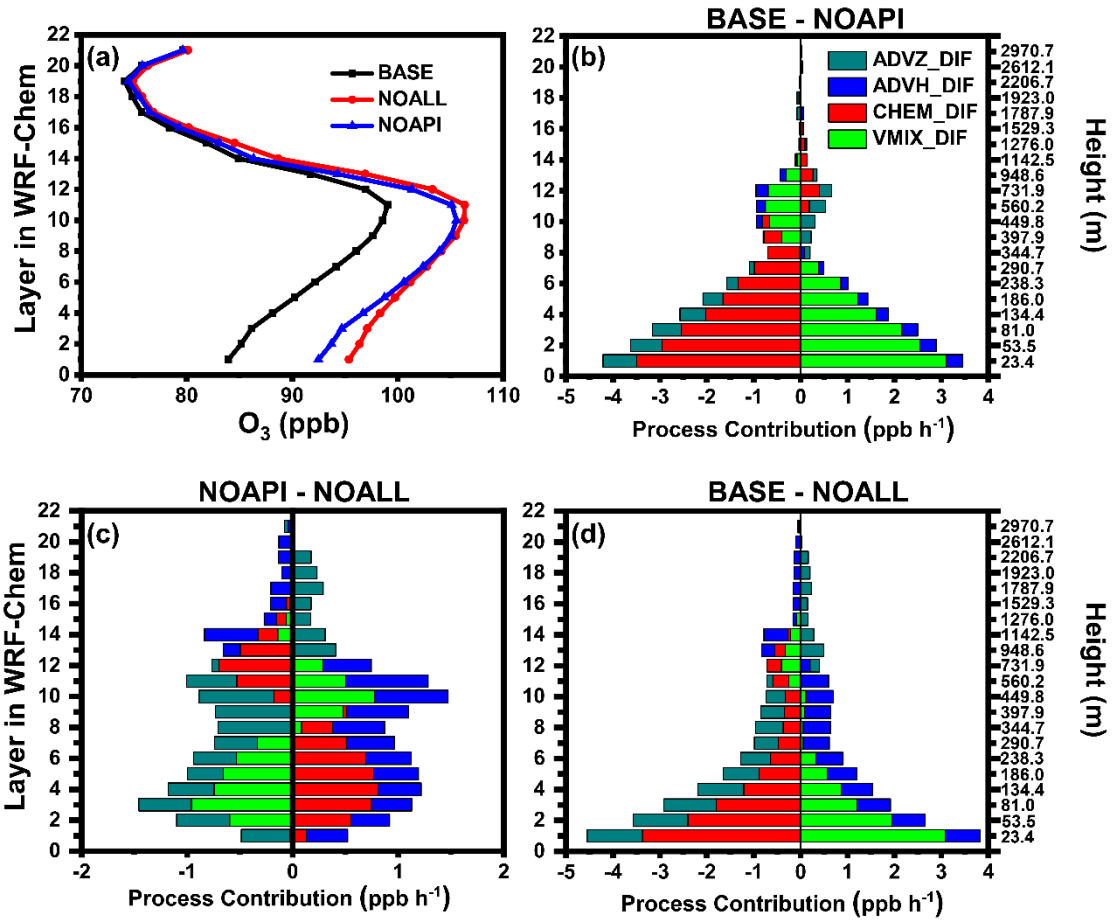


Figure 7. (a) Diurnal variations of simulated surface O_3 concentrations in BASE (black), NOAPI (blue), and NOALL (red) cases over CAPAs. (b) The hourly surface O_3 changes induced by each physical/chemical process using the IPR analysis method in BASE case. (c-e) Changes in hourly surface O_3 process contributions caused by API (BASE minus NOAPI), ARF (NOAPI minus NOALL), and ALL (BASE minus NOALL) over CAPAs during the daytime (08:00-17:00 LST) from 28 July to 3 August 2014. The black lines with squares denote the net contribution of all processes (NET, defined as VMIX+CHEM+ADV). Differences of each process contribution are denoted as VMIX_DIF, CHEM_DIF, ADV_DIF, and NET_DIF.


(a)
BASE
NOALL
NOAPI
Layer in WRF-Chem
O3 (ppb)
BASE - NOAPI
(b)
ADVZ_DIF
ADVH_DIF
CHEM_DIF
VMIX_DIF
Process Contribution (ppb h-1)
Height (m)
NOAPI - NOALL
(c)
BASE - NOALL
(d)

1

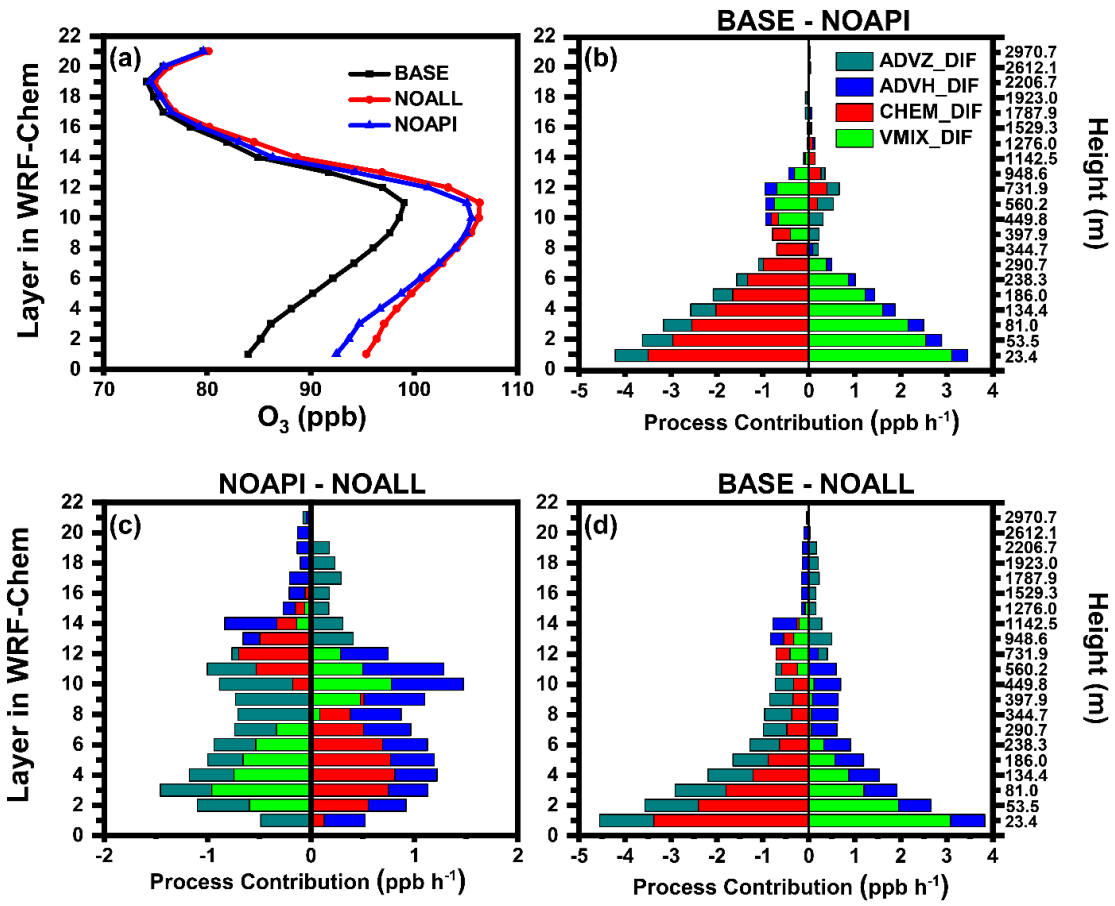


Figure 8. (a) Vertical profiles of simulated O_3 concentrations in BASE (black), NOAPI (blue), and NOALL (red) cases over CAPAs. (b-d) Changes in O_3 budget due to API, ARF, and ALL over CAPAs during the daytime (08:00-17:00 LST) from 28 July to 3 August 2014. Differences of each process contribution are denoted by ADVZ_DIF, ADVH_DIF, CHEM_DIF, and VMIX_DIF.