

1



# <sup>2</sup> Quantifying NO<sub>x</sub> emissions in Egypt using TROPOMI observations

<sup>3</sup> Anthony Rey-Pommier<sup>1</sup>, Frédéric Chevallier<sup>1</sup>, Philippe Ciais<sup>1</sup>, Grégoire Broquet<sup>1</sup>, Theodoros

<sup>4</sup> Christoudias<sup>2</sup>, Jonilda Kushta<sup>2</sup>, Didier Hauglustaine<sup>1</sup> and Jean Sciare<sup>2</sup>.

<sup>5</sup> Laboratoire des Sciences du Climat et de l'Environnement, LSCE/IPSL, CEA-CNRS-UVSQ, Université Paris-Saclay,

6 91190 Gif-sur-Yvette, France

 $_{7}$   $^{2}$  The Cyprus Institute, Climate and Atmosphere Research Center, 2121 Nicosia, Cyprus

**Correspondence:** Anthony Rey-Pommier (anthony.rey-pommier@lsce.ipsl.fr)

Abstract. Urban areas and industrial facilities, which concentrate most human activity and industrial production, are major sources of air pollutants, with serious implications for human health and global climate. For most of these 10 pollutants, emission inventories are often highly uncertain, especially in developing countries. Spaceborne observations 11 from the TROPOMI instrument, onboard the Sentinel-5 Precursor satellite, are used to measure nitrogen dioxide (NO<sub>2</sub>) 12 slant column densities with a high spatial resolution. Here, we use two years of TROPOMI retrievals to map nitrogen 13 oxides  $(NO_x = NO + NO_2)$  emissions in Egypt with a top-down model based on the continuity equation in steady 14 state. Emissions are expressed as the sum of a transport term and a sink term representing the three-body reaction 15 comprising  $NO_2$  and OH. This sink term requires information on the lifetime of  $NO_2$ , which is calculated with the 16 use of CAMS near-real-time temperature and hydroxyl radical (OH) concentration fields. The applicability of the 17 OH concentration field is evaluated by comparing the lifetime it provides with the lifetime inferred from the fitting of 18  $NO_2$  line density profiles with an exponentially modified Gaussian function. This comparison, which is conducted for 19 39 samples of NO<sub>2</sub> patterns above the city of Riyadh, provides information on the reliability of the CAMS near-real-20 time OH concentration fields; It also provides the location of the most appropriate vertical level to represent typical 21 pollution sources in industrial areas and megacities in the Middle East. In Egypt, total derived emissions of  $NO_x$  are 22 dominated by the sink term. However, they can be locally dominated by wind transport, especially along the Nile 23 where human activities are concentrated. Megacities and industrial regions clearly appear as the largest sources of 24  $NO_x$  emissions in the country. Our top-down model produces emissions whose annual variability is consistent with the 25 national electricity consumption. It is also able to detect lower emissions on Fridays, which are inherent to the social 26 norm of the country, and to quantify the drop in emissions due to the COVID-19 pandemic. Overall, our indications of 27 NO<sub>x</sub> emissions for Egypt are found to be 25.0% higher than the CAMS-GLOB-ANT v4.2 inventory, but significantly 28 differ in terms of seasonality. 29

30

#### 31 1 Introduction

Economic growth in developing countries has led to a strong increase of urban air pollution (Baklanov et al., 2016 [1]). 32 Among the different pollutants, nitrogen oxides are key species. They are generally the products of fuel combustion, 33 such as the combustion of hydrocarbons in the air at high temperature. The main sources of these compounds are 34 therefore vehicle engines, but also heavy industrial facilities such as power plants, metal smelters and cement plants. 35 Their accumulation in the lowest layers of the troposphere contributes to the formation of smog and acid rain (Singh 36 et al., 2007 [2]). They also have a significant effect on human health, as they can cause various respiratory diseases (EPA, US., 2016 [3]). To deal with these phenomena, national and regional governments generally enact a series of air 38 pollution control strategies, which typically take the form of bans on certain polluting technologies, with the aim of reducing the concentration of pollutants at the local level to targets that must be achieved within a given timeframe. 40





These strategies, which also help driving technological innovation, have had a significant effect in Europe (Crippa et al., 2016 [4]).

In Egypt, population growth, urbanisation, socio-economic development and the associated increase in the vehicle 43 fleet led to a major degradation of air quality in the last decades, especially in highly populated areas such as Greater 44 Cairo and the Nile Delta (Abou El-Magd et al., 2020 [5]), which gather the majority of the population. The Ministry 45 of State for the Environment has thus initiated new policies that aim to reduce pollution levels throughout the country, 46 through technical mitigation of emissions, emission standards for vehicles and intersectoral collaboration (UNEP, 2015 47 [6]). However, most developing countries, such as Egypt, lack the local infrastructure to access detailed information 48 on technical factors such as energy consumption, fuel type and technology, leading to discrepancies in inventories (Xue 49 et al., 2012 [7]). As a consequence, the monitoring of emissions, which is important to evaluate the effects of the air 50 pollution control policies, is of limited reliability. 51 To overcome the uncertainties in the emission inventories, the use of independent observations systems is becoming

52 increasingly prevalent. In this study, we investigate the use of satellite remote sensing of atmospheric concentrations to 53 improve the quantification of  $NO_x$  emissions in Egypt. Spectrally resolved satellite measurements of solar backscattered 54 radiation enable the quantification of  $NO_2$  and other trace gases absorbing in the UV-Visible spectral range based 55 on their characteristic spectral absorption patterns. Tropospheric vertical column densities, i.e. vertically integrated 56  $NO_2$  concentrations in the troposphere, have been providing information on the spatial distribution of tropospheric 57  $NO_2$  at global scale for nearly 30 years, allowing the identification of different sources of  $NO_x$  and the quantification 58 of the associated emissions (Leue et al., 2001 [8]; Martin et al., 2003 [9]; Mijling and van der A, 2012 [10]; de Foy et 59 al., 2015 [11]; Goldberg et al., 2019 [12]; Beirle et al., 2019 [13]; Lorente et al., 2019 [14]; Lange et al., 2021 [15]). In 60 October 2017, the Sentinel-5 Precursor satellite was launched. Its main instrument is the TROPOspheric Monitoring 61 Instrument (TROPOMI), which provides tropospheric  $NO_2$  column densities at a high spatial resolution, with a large 62 swath width and with a daily frequency (Veefkind et al., 2012 [16]). By applying the steady-state continuity equation 63 (Beirle et al., 2019 [13]; Lama et al., [17], 2020), it is possible to build a top-down model that directly quantifies  $NO_x$ 64 emissions from these NO<sub>2</sub> column densities, provided that some key parameters (wind, temperature, hydroxyl radical 65 concentration and concentration ratio between  $NO_x$  and  $NO_2$ ) are correctly estimated. This model is used to quantify 66 the anthropogenic  $NO_x$  emissions in Egypt for a 2-year period, from November 2018 to November 2020. 67

This paper is organised as follows: Section 2 provides a description of the datasets used in this study. Section 3 68 explains the build-up and the limits of the top-down model used to quantify  $NO_x$  emissions in Egypt. It also presents 69 a method for validating the parameters of the model by using  $NO_2$  line density profiles over Riyadh, Saudi Arabia. 70 Section 4 presents the analysis of this validation method. It presents the location of the main  $NO_x$  sources in Egypt 71 and evaluates the vertical sensitivity of the model. It also assesses the ability of the model to show less human activity 72 on Fridays and during the lockdown that took place during the COVID-19 pandemic. It finally confronts the inferred 73 emissions with different inventories in terms of amplitude and seasonality. Section 5 presents our conclusion and 74 general remarks. 75

#### 76 2 Instrumentation and data

#### 77 2.1 TROPOMI NO<sub>2</sub> retrievals

The TROPOspheric Atmosphere Monitoring Instrument (TROPOMI), onboard the European Space Agency's (ESA) 78 Sentinel-5 Precursor (S-5P) satellite, provides measurements for atmospheric composition. TROPOMI is a spectrom-79 eter observing wavelengths in the infrared, visible and ultraviolet light at around 13:30 local time. The UV-Visible 80 spectral band at 405-465 nm is used to retrieve NO<sub>2</sub>. Other spectral bands are used to observe methane, formaldehyde, 81 sulphur dioxide, carbon monoxide and ozone, as well as aerosols and cloud physical properties. The very high spatial 82 resolution offered by TROPOMI (originally  $3.5 \times 7 \text{ km}^2$  at nadir, improved to  $3.5 \times 5.5 \text{ km}^2$  since 6 August 2019) 83 provides unprecedented information on tropospheric NO<sub>2</sub>. Its large swath width ( $\sim 2600$  km) enables to construct 84 NO<sub>2</sub> images on large spatial scales. Those images greatly improve the potential for detecting highly localised pollu-85 tion plumes above the ground, identifying small-scale emission sources but also estimating emissions from megacities, 86 industrial facilities and biomass burning. We use TROPOMI NO<sub>2</sub> retrievals from November 2018 to November 2020 87 over the Middle East and Eastern Mediterranean region, and more specifically over Egypt and the city of Riyadh, Saudi Arabia. The arid climate of this region offers a large number of clear-sky days throughout the year, but also 89 the presence of large plumes of pollutants due to a large human concentration along rivers and around megacities, which allows us to observe high NO<sub>2</sub> concentrations in the region with a high signal-to-noise ratio. To facilitate data 91





- <sup>92</sup> filtering, TROPOMI products provide a quality assurance value  $q_a$ . This value ranges from 0 (no data) to 1 (high-
- $q_a$  quality data). For our analysis of concentrations, we selected NO<sub>2</sub> retrievals with  $q_a$  values equal to or greater than
- 94 0.75, which systematically correspond to clear-sky conditions (Eskes and Eichmann, 2019 [18]). TROPOMI soundings
- are gridded for this study at a spatial resolution of  $0.1^{\circ} \times 0.1^{\circ}$  with daily coverage.

#### 96 2.2 Wind data

- $\mathbf{v}_{\tau}$  The horizontal wind information  $\mathbf{w} = (u, v)$  is taken from the European Centre for Medium-Range Weather Forecasts
- (ECMWF) ERA5 data archive (fifth generation of atmospheric reanalyses) at a horizontal resolution of  $0.25^{\circ} \times 0.25^{\circ}$
- on 37 pressure levels (Hersbach et al., 2020 [19]). The hourly values have been linearly interpolated to the TROPOMI
- orbit timestamp and re-gridded with a  $0.1^{\circ} \times 0.1^{\circ}$  resolution.

## <sup>101</sup> 2.3 CAMS real-time fields

The Copernicus Atmospheric Monitoring Service (CAMS) global near-real-time service provides analyses and forecasts 102 for reactive gases, greenhouse gases and aerosols on 25 pressure levels with a horizontal resolution of  $0.4^{\circ} \times 0.4^{\circ}$  and 103 a temporal resolution of 3 hours (Huijnen et al., 2016 [20]). For the calculation of  $NO_x$  emissions from TROPOMI 104 observations, we use CAMS concentration fields of nitrogen oxides (NO and NO<sub>2</sub>) and hydroxyl radical (OH). We also 105 use CAMS temperature field T. NO and NO<sub>2</sub> concentrations are used to account for chemical processes that take place 106 in polluted air. Anthropogenic activities produce mainly NO, which is transformed into  $NO_2$  by reaction with ozone 107  $O_3$ . NO<sub>2</sub> is then photolyzed during the day, reforming NO and producing an oxygen atom O (Seinfeld, 1989 [21]). 108 This photochemical equilibrium between NO and  $NO_2$  can be highlighted with the  $NO_x:NO_2$  concentration ratio, 109 whose value depends on local conditions, allowing to perform a conversion from  $NO_2$  production to  $NO_x$  emissions. 110 The reason for the use of OH is different. OH is the main oxidant that controls the ability of the atmosphere to 111 remove pollutants such as NO<sub>2</sub> (Logan et al., 1981 [22]). It is mainly produced during daylight hours by interaction 112 between water and atomic oxygen produced by ozone dissociation (Levy, 1971 [23]). In air that is directly influenced 113 by pollution, the second source of OH is due to a reaction between NO and  $HO_2$ . This reaction, referred to as the  $NO_x$ 114 recycling mechanism, illustrate the nonlinear dependence of the OH concentration on NO<sub>2</sub> (Valin et al., 2011 [24]: 115 Lelieveld et al., 2016 [25]). Since the OH lifetime is typically of less than a second, its concentration in the troposphere 116 is very low and difficult to measure. As a consequence, global analyses, which estimate OH concentrations from other 117 variable species (Li et al., 2018 [26]; Wolfe et al., 2019 [27]), provide a natural representation for OH concentrations, 118 but with high associated uncertainties. Therefore, the CAMS OH concentrations are used here to account for the 119  $NO_2$  oxidation to form nitric acid (HNO<sub>3</sub>), whose representation is explained in Section 3.1. Finally, the temperature 120 field is used to control variations in the kinetic parameters (Burkholder et al., 2020 [28]). The hourly values are also 121 linearly interpolated to the TROPOMI orbit timestamp and re-gridded with a  $0.1^{\circ} \times 0.1^{\circ}$  resolution. 122

#### 123 2.4 Calculation of urban enhancements

Detecting traces of anthropogenic emissions in TROPOMI NO<sub>2</sub> images is not a straightforward process. The NO<sub>2</sub> 124 signal from a sparsely populated area or a small industrial plant may be covered by numerical noise or by the signal 125 generated by natural sources such as lightning (Boersma et al., 2005 [29]) and soil emissions (Yienger et al., 1995 [30]). 126 It is therefore necessary to remove the natural part of the atmospheric signal from the detected  $NO_2$  enhancement. 127 With an atmospheric lifetime of about a few hours, the presence of  $NO_2$  is relatively short. Consequently, the majority 128 of  $NO_2$  is not transported far downwind from its sources. Thus, near-surface  $NO_2$  concentrations are generally high 129 over industrial facilities and densely populated areas that need to be identified. Because Egypt's population is almost 130 entirely located along the Nile River and its delta, the study of  $NO_x$  emissions in this country cannot therefore be 131 reduced to the study of a small number of point sources, as it would be the case for several other parts of the region, 132 and must be carried out in the form of a mapping of the country. Further explanation is provided in Section 3.3. 133 To identify urban or industrial areas in Egypt, we use the Socioeconomic Data and Applications Center (SEDAC) 134 GRUMP (Global Rural-Urban Mapping Project) data archive, which comprises eight global datasets, including a 135 population density grid provided at a resolution of 30 arc seconds, with population estimates normalised for the year 136 2000 (CIESIN, 2019 [31]). We combine this database with field data giving the location of industrial facilities from 137 energy-intensive industries in the region (data have been retrieved from Global Energy Observatory for oil and gas-138 fired power plants, from IndustryAbout for aluminium, steel and iron plants, from the work of Elvidge et al., 2016 139 [32] for flaring sites, and from the work of Steven J. Davis and Dan Tong for cement plants, links are given at the end 140 of this article). 141







Figure 1: (left) Part of Egypt centered on Nile river. Pink cells represent locations with an average human density above 100 hab.km<sup>-2</sup>, brown cells represent locations with industrial facilities outside cities, and yellow cells represent locations in the vicinity of pink and brown cells. Green cells represent areas which do not correspond to any of the three criteria, considered to be void of human activity. Five larges cities in the country and the industrial area of Ain Sokhna are denoted with stars. (right) TROPOMI observation of NO<sub>2</sub> column densities above Nile valley on 3 January 2019. White pixels correspond to areas with low quality data ( $q_a < 0.75$ ) or no data.

These datasets are used to remove the non-anthropogenic part of the  $NO_x$  emissions signal. We conduct this removal 142 by subtracting the mean emissions over desert and rural areas from the mean emissions over urban and industrial 143 areas. In order to perform this distinction between two types of areas, our study is carried out using a grid with 144 a resolution of  $0.1^{\circ} \times 0.1^{\circ}$  characterised by two types of cells. A cell is considered "urban" if it has a population 145 density higher than a threshold of 100 hab  $\mathrm{km}^{-2}$ , or if its centre is close to an industrial facility. Otherwise, the cell 146 is considered "rural". In order to avoid any smearing that would correspond to abnormally high emissions outside 147 urban areas (which can happen if the wind is poorly estimated), transition cells (in the immediate vicinity of the 148 mentioned urban cells) are also considered to be urban. Figure 1 shows the distinction between urban and rural cells 149 on our domain in Egypt that lies between parallels 23°N and 32°N and meridians 29°E and 34°E. Most urban cells 150 are located in the Nile area. Some urban cells are also found on the coast or in isolated parts in the desert. They 151 correspond to remote industrial facilities, including major flaring sites, or sparsely populated industrial centres such as 152 Ain Sokhna's industrial area. The domain comprises  $n_{rur} = 3692$  rural cells and  $n_{urb} = 949$  urban cells. The resulting 153 grid is conveniently used as a mask for the urban enhancement, whose calculation is explained in Section 3.4. 154

#### 155 2.5 Emission inventories

We compare TROPOMI-derived  $NO_x$  emissions to the Emissions Database for Global Atmospheric Research 156 (EDGARv5.0) for 2020 and the CAMS global anthropogenic emissions (CAMS-GLOB-ANT v4.2) inventory released 157 in 2020. Both datasets provide  $0.1^{\circ} \times 0.1^{\circ}$  gridded emissions for different sectors on a monthly basis. EDGARv5.0 158 (Crippa et al., 2019 [33]) is based on activity data (population, energy production, fossil fuel extraction, industrial pro-159 cesses, agricultural statistics, etc.) derived from the International Energy Agency (IEA) and the Food and Agriculture 160 Organization (FAO), corresponding emission factors, national and regional information on technology mix data and 161 end-of-pipe measurements. The inventory covers the years 1970-2015. CAMS-GLOB-ANT v4.2 (Granier et al., 2019 162 [34]) is developed within the framework of the Copernicus Atmospheric Monitoring Service. Its NO<sub>x</sub> emissions are 163 based on various existing sectors in the EDGARv4.3.2 emissions from 2000-2012 and are extrapolated to the current 164 year using 2011-2014 sector-based trends from the Community Emissions Data System (CEDS) inventory (Hoesly 165 et al., 2018 [35]). From one inventory to another, the names of the sectors may vary. In EDGARv5.0 and CAMS-166 GLOB-ANT\_v4.2, the emissions for a given country are derived from the type of technologies used, the dependence 16 of emission factors on fuel type, combustion conditions, as well as activity data and low resolution emission factors 168 (Janssens-Maenhout et al., 2019 [36]). 169



173

207

208



#### 3 Method 170

#### 3.1 Calculation of NO<sub>2</sub> production from TROPOMI observations 171

As a first step, we use TROPOMI's tropospheric NO<sub>2</sub> columns  $\Omega_{NO_2}$  to derive top-down NO<sub>2</sub> production maps. Using 172

the horizontal wind  $\mathbf{w}$ , the NO<sub>2</sub> flux is given as  $\Omega_{NO_2}\mathbf{w}$ . The divergence of this flux is added to the local time derivative  $\frac{\partial \Omega_{\text{NO}_2}}{\partial t}$  to balance NO<sub>2</sub> sources  $e_{\text{NO}_2}$  and sinks  $s_{\text{NO}_2}$  according to the continuity equation: 174

$$\frac{\partial \Omega_{\rm NO_2}}{\partial t} + \operatorname{div}(\Omega_{\rm NO_2} \mathbf{w}) = e_{\rm NO_2} - s_{\rm NO_2} \tag{1}$$

In steady state, the time derivative disappears and the mass balance is reduced to three terms. The sources of  $NO_2$ 175 can be estimated by taking the combined effect of atmospheric transport losses and the different sinks. For the 176 transport term, we calculate numerical derivatives with a fourth-order central-finite difference formula at each point 177 of the domain. For the sink term, since the local overpass time of TROPOMI occurs is in the middle of the day,  $NO_2$ 178 losses are largely dominated by the three-body reaction involving NO<sub>2</sub> and OH (Seinfeld, 1989 [21]). Two channels 179 have been identified for this reaction (Burkholder et al., 2020 [28]), leading to the production of nitric acid HNO<sub>3</sub> and 180 pernitrous acid HOONO: 181

- $NO_2 + OH + M \rightarrow HNO_3 + M$ 182
- $NO_2 + OH + M \rightarrow HOONO + M$ 183

No corrections are made to the TROPOMI observations. Slant column densities are used as vertical densities. This 184 use amounts to neglect the air mass factor, which is a source of structural uncertainty in  $NO_2$  retrievals (Boersma et 185 al., 2004 [37]; Lorente et al., 2017 [38]). In polluted environments, this source of uncertainty becomes non-negligible. 186 Rather than calculating the air mass factor, we take this factor into account in the final uncertainty estimates. The 187 corresponding uncertainty level is given in Section 4.7. For the OH concentrations that are considered in this region 188  $(1-30 \times 10^6 \text{ molecules.cm}^{-3})$ , the reactions above follow first order kinetics. The total sink term can therefore be calculated as  $s_{NO_2} = \Omega_{NO_2}/\tau$  with: 189 190

$$\tau = \frac{1}{k_{mean}(T, [M]) \cdot [OH]}$$
(2)

 $\tau$  appears here as the characteristic mixed lifetime of NO<sub>2</sub> in the atmosphere. The value of the reaction rate  $k_{mean}$ 191 characterises the reactions between  $NO_2$  and OH and depends on atmospheric conditions (temperature T and total air 192 concentration [M]). We calculate the value of  $k_{mean}$  using a temperature-dependent analytical formula for different 193 pressure ranges (Burkholder et al., 2020 [28]). Note that although this reaction rate accounts for both reactions with OH, the second channel is minor and cannot be considered as a true  $NO_x$  sink, HOONO being rapidly decomposed back to NO<sub>2</sub> and OH in the lower troposphere (Sander et al., 2011 [39]). The value of  $k_{mean}$  therefore represents the total loss of NO<sub>2</sub> due to OH and cannot be used to infer HNO<sub>3</sub> and HOONO production. Thus, the NO<sub>2</sub> production can be calculated as the sum of a transport term and a sink term: 198

$$e_{\mathrm{NO}_2} = \operatorname{div}(\Omega_{\mathrm{NO}_2} \mathbf{w}) + \Omega_{\mathrm{NO}_2} / \tau \tag{3}$$

The treatment for  $NO_x$  removal is simplified here. In reality,  $NO_x$  concentrations are influenced by other sinks. At 199 the global scale, the other important sinks are: 200

- NO<sub>2</sub> deposition through the leaf stomata of vegetation (Delaria et al., 2020 [40]); 201

- NO<sub>x</sub> oxidation by organic radicals to produce of alkyl and multifunctional nitrates (Sobanski et al., 2017 [41]; 202 Romer Present et al., 2020 [42]); 203

- NO reaction with HO<sub>2</sub> to produce nitric acid (Butkovskaya et al., 2005 [43]), whose yield is strongly enhanced 204 in presence of water vapour (Butkovskaya et al., 2009 [44]); 205

- NO conversion to NO<sub>3</sub>, the latter being in thermal equilibrium with NO<sub>2</sub> and  $N_2O_5$ ; 206
  - NO<sub>2</sub> reversible reaction with peroxyacetyl radical to produce peroxyacetylnitrate (Moxim et al., 1996 [45]);
  - NO<sub>2</sub> uptake onto black carbon particles (Longfellow et al., 1999 [46]).

Stavrakou et al., 2013 [47] has shown that the reaction between  $NO_2$  and OH forming HNO<sub>3</sub> accounted for 26 to 209 64% of total NO<sub>x</sub> emissions at the global scale. However, the features of the climate in Egypt during daytime hinder 210 many other sinks to significantly take place. Losses due to deposition and the formation of alkyl and multi-functional 211 212 nitrates are thus considered insignificant in Egypt where the forest cover is totally negligible. We also neglect the reaction between NO and  $HO_2$  as the corresponding reaction rate is weakened by a factor 3 to 8 in dry conditions 213 (Butkovskaya et al., 2005 [43]). The formation of  $NO_3$  and  $N_2O_5$  via heterogeneous processes, which has a significant 214 contribution during nighttime in the Mediterranean region (Friedrich et al., 2021 [48]), is neglected at 13:30 when OH 215





<sup>216</sup> is almost at its daily maximum. Similarly, the production of PAN, which peaks in late afternoon and early evening <sup>217</sup> (Seinfeld, 1989), is neglected at 13:30. Finally, the uptake of  $NO_2$  onto black carbon is of limited amount in the <sup>218</sup> Mediterranean region (Friedrich et al., 2021 [48]). All these processes being neglected, the reaction between  $NO_2$  and <sup>219</sup> OH is the only sink that is considered in our calculations to provide a reliable indication of  $NO_x$  emissions.

#### <sup>220</sup> 3.2 Interpolation to daily average emissions

All parameters are evaluated at 13:30 local time, which means  $NO_2$  production is calculated for the same moment. In Egypt, the maximum and minimum electricity consumption are reached around 20:00 and 6:00 local time respectively,

and inter-daily consumption differences have been weakened by the increasing sales of air conditioning and ventilation

systems in the past decades (Attia et al., 2012 [49]). The national average daily load profiles provided by the National

225 Egyptian Electricity Holding Company show that the mean daily electricity consumption corresponds approximately

to the consumption at 13:30 in the country (EEHC, 2021 [50]). The difference between the two quantities being very

 $_{227}$  small both in summer (about +2 to -3%) and winter (about -2 to -6%), we neglect this difference and consider our

<sup>228</sup> inferred emissions as representative of the average activity in Egypt at any time of the year.

#### 229 3.3 Validation of CAMS OH concentration using line density calculations for Riyadh

When the transport term is integrated over large spatial scales, it cancels out due to the mass balance in the continuity 230 equation between  $NO_2$  sources and  $NO_2$  sinks. At first order, the integration of the inferred emissions over the whole 231 domain (of about  $450,000 \text{ km}^2$ ) thus reflects chemical losses of the sink term. In this term, the NO<sub>2</sub> lifetime calculation 232 involves the reaction rate  $k_{mean}$ , whose annual variability is low due to small changes in Egyptian midday temperatures 233 throughout the year, and OH concentration, whose annual variability is highly marked. In Egypt, tropospheric OH 234 concentrations, which are strongly correlated with solar ultraviolet radiation (Rohrer and Berresheim, 2006 [51]) and 235 NO<sub>x</sub> emissions, are higher in summer than in winter. To ensure an adequate representation of the OH field by CAMS 236 data, we select a large number of TROPOMI images characterised by a homogeneous wind field, in which we calculate 237 the NO<sub>2</sub> lifetime according to Equation (2), where [OH] corresponds to the near-real-time CAMS data and  $k_{mean}$  is 238 calculated with the formula from Burkholer et al., 2020 [28]. We compare this value with the lifetime determined by a 239 method initially developed by Beirle et al., 2011 [52], and expanded by Valin et al., 2013 [53] by introducing a rotation 240 of the image to refine the chemical lifetime. This method consists in fitting an exponentially modified Gaussian 241 function (EMG) to  $NO_2$  line density profiles. These profiles correspond to the  $NO_2$  mean value along the mean wind 242 direction in the pollution pattern and centered around the source. They are obtained by rotating TROPOMI images 243 in the mean wind direction and averaging the values of the nearest columns in a 100  $\rm km^2$  area. Line density profiles 244 are generated on a distance of 300 km. An example is given in Figure 3. Within the average profile, the NO<sub>2</sub> burden 245 and lifetime can be derived from the parameters that describe the best statistical fit. The EMG model is expressed 246 as follows (Lange et al., 2021 [15]): 247

$$\langle \Omega_{\rm NO_2} \rangle (x|B, A, x_0, \mu, \sigma) = B + \frac{A}{2x_0} \exp\left(\frac{\mu - x}{x_0} + \frac{\sigma^2}{2x_0^2}\right) \operatorname{erfc}\left(-\frac{1}{\sqrt{2}}\left(\frac{x - \mu}{\sigma} - \frac{\sigma}{x_0}\right)\right) \tag{4}$$

Here, x is the distance in the downwind-upwind direction, B is the NO<sub>2</sub> background, A is the total number of NO<sub>2</sub> molecules observed in the vicinity of the point source,  $x_0$  is the e-folding distance downwind, representing the exponential length scale of NO<sub>2</sub> decay,  $\mu$  is the location of the apparent source relative to the centre of the point source, and  $\sigma$  is the standard deviation of the Gaussian function, representing the length scale of Gaussian smoothing. Using a non-linear least squares fit, we estimate the five unknown parameters: A, B,  $x_0$ ,  $\mu$  and  $\sigma$ . Using the mean zonal wind module  $\sqrt{\mathbf{w}^2}$  of the NO<sub>2</sub> line density domain, the mean effective NO<sub>2</sub> lifetime  $\tau_{\text{fit}}$  can be estimated from the fitted parameters:

$$\tau_{\rm fit} = \frac{x_0}{\langle \sqrt{\mathbf{w}^2} \rangle} \tag{5}$$

The geography of Egypt does not suit the method described here. The Egyptian population is contiguously concentrated along the Nile, which makes it difficult to define point sources isolated from human activity. Furthermore, large isolated cities such as Alexandria or Suez are too close to the coast for the wind to be considered homogeneous. We therefore use the nearby city of Riyadh, Saudi Arabia (24.684°N, 46.720°E) to perform the comparison between the CAMS-induced lifetime and the fit-induced lifetime. Riyadh is particularly suitable for several reasons. Firstly, Riyadh is a city within the latitudinal extend of Egypt (1600 km from Cairo), has a climate which is similar to the typical Egyptian climate. Second, NO<sub>2</sub> tropospheric columns over Riyadh are high (~ 9 × 10<sup>15</sup> molecules.cm<sup>-2</sup>),





202 leading to retrievals with a high signal-to-noise ratio. Third, Riyadh is far from the coast, and its flat terrain makes 203 the surrounding wind fields rather homogeneous during most of the year.

As the fitting algorithm is very sensitive to any disturbance that might be induced by  $NO_2$  production from other 264 point sources, it is necessary to identify heavy industrial facilities in the area. As Riyadh is also an industrial area, 265 many power plants are located close to the city centre. Figure 2 shows the location of the most important emitters in 266 the region, which include five gas-fired power plants (PP7, PP8, PP9, PP10 and PP14), one oil-fired power plant (PP4) 267 and one cement plant (CP). The five gas power plants, with a total capacity of more than 10 GWe, are located in the 268 periphery of the city. These power plants are sufficiently far away from the city centre for TROPOMI to distinguish 269 their own emissions from those of Riyadh's centre itself with a resolution of  $0.1^{\circ} \times 0.1^{\circ}$ , which is not the case for CP 270 and PP4 which are located in the city centre. It is therefore appropriate to restrict the study of  $NO_2$  patterns over 271 Riyadh to days for which the emissions from the city centre and those from the gas power plants do not mix. This is 272 the case when the wind blows steadily and homogeneously in a north-south direction. Within about 150 km around 273 the city centre, we thus calculate the average zonal wind given by ERA5 and consider the observation as reliable if the 274 mean angle  $\langle \theta \rangle$  of the observations deviates by less than 40° from the north or the south, with a standard deviation 275  $\sigma_{\theta}$  of less than 36°. This condition generally leads to a selection of observations with large wind speeds, low winds 276 speeds being often associated with more variable directions. This ensures the  $NO_2$  decay to be dominated by chemical 277 removal and not by the variability of the wind (Valin et al., 2013 [53]). Finally, we select observations with clear-sky 27 conditions downstream of the flow (with at least 80% downstream cells with  $q_a > 0.75$ ). 279



Figure 2: Map of Riyadh's city centre with its surrounding power plants (PP4, 7, 8, 9, 10 and 14) and cement plant (CP). The map has been extracted from © Google Maps.

Our  $0.1^{\circ} \times 0.1^{\circ}$  gridding ensures that retrieved lifetimes are governed by physical decay of NO<sub>2</sub> and not an artifact of 280 the spatial resolution (Valin et al., 2011 [24]). The fitting procedure is very sensitive to the wind direction. Instead of 281 manually correcting the ERA5 wind field for individual NO<sub>2</sub> patterns, we use another procedure: the curve fitting is 282 performed for every sample with three different rotation angles, corresponding to the wind direction with a correction 283 of -10, 0 or 10 degrees. A record is kept if one of these three fits leads to a determination correlation with the 284 corresponding  $NO_2$  line density whose coefficient is greater than 0.97. Among the remaining samples, we keep those 28 with a value of  $\tau_{\rm fit}$  greater than 1.0 hr (considered sufficiently high to be relevant). An example of curve fitting is 280 given in Figure 3. 287

The phenomena under study take place in the planetary boundary layer (PBL), which in this region has a midday 288 height of about 2 km (Filioglou et al., 2020 [54]). TROPOMI observations only provide information on the total 289 NO<sub>2</sub> content of the tropospheric column, without providing information on the vertical distribution of concentrations. 290 Extracting emissions from concentrations therefore requires a choice on the height at which wind, temperature and 291 OH data are taken. Lama et al., 2020 [17] and Lorente et al., 2019 [14] conducted similar studies using the boundary 292 layer average wind, while Beirle et al., 2019 [13] chose a vertical level of about 450 m above ground. Because vertical 293 transport of NO<sub>x</sub>, which is emitted mainly from combustion engines and industrial stacks, is generally minor compared 294 to horizontal transport, NO<sub>x</sub> is confined to the first few hundred metres above ground level. Using PBL-averaged data 295 poses a problem of consistency as wind, temperature and OH concentration values vary significantly within the PBL. 296 As a consequence, we compare the CAMS-induced lifetime  $\tau$  and the fit-induced lifetime  $\tau_{\rm fit}$  using the parameters (w, 29 [OH] and T) at two different vertical levels: a medium level A at 925 hPa (about 770 m above ground level), and 298





a bottom level  $\mathcal{B}$  at 987.5 hPa (about 210 m). These levels are interpolated from four and two ECMWF or CAMS consecutive pressure levels respectively (1000-850 hPa for level  $\mathcal{A}$  and 1000-975 hPa for level level  $\mathcal{B}$ ). Most urban cells having an altitude between 0 and 150 m, the corresponding pressure variations are small (up to ~ 16 hPa), which allows us to neglect the effects of topography on the position of pressure levels. Figure 4 sums up the selection method for the comparison of methods.



Figure 3: Estimation of the NO<sub>2</sub> lifetime from a pattern above Riyadh on 03/08/2020: (left) NO<sub>2</sub> plume rotated with its wind direction around the source (cross) to an upwind-downwind pattern. Grey boxes centered around black points indicate the spatial extent in the spatial integration of NO<sub>2</sub> columns to obtain the NO<sub>2</sub> line density. Values of cardinal points are noted in black. (right) Line densities (points) representing the downwind evolution of NO<sub>2</sub> as function of the distance to Riyadh city centre, and the corresponding fit according to the exponentially modified Gaussian function (dashed line).



Figure 4: Selection method for NO<sub>2</sub> patterns over Riyadh. Datasets (yellow-orange) are used to calculate the quantities (light green) that are submitted to different tests (grey). 731 patterns are progressively conserved (green arrows) or rejected (red arrows). At each stage, the number of conserved or rejected patterns are noted within brackets (the value is only given for calculations performed at level  $\mathcal{B}$ ). This selection process compares the lifetimes estimated by the EMG function fitting with TROPOMI line density profiles to the lifetimes calculated according to Equation (2) with CAMS data.

#### <sup>304</sup> 3.4 Calculation of anthropogenic NO<sub>x</sub> emissions and comparison with inventories

We calculate  $NO_x$  emissions on the entire domain from  $NO_2$  production by using CAMS NO and  $NO_2$  concentrations. These quantities are not intended to replace TROPOMI observations; they are used to apply the concentration ratio





 $[NO_x]/[NO_2] = ([NO]+[NO_2])/[NO_2]$  to account for the conversion of NO<sub>2</sub> to NO and vice versa. As NO concentrations 307 in urban areas generally range from 2 to 10 ppb, the characteristic stabilization time of this ratio ranges from 4 to 20 308 minutes (Graedel et al., 1976 [55]; Seinfeld and Pandis, 2006 [56]). This time being lower than the order of magnitude 309 of the inter-mesh transport time (about 30 min considering the resolution used and the mean wind module in the 310 region), we can reasonably neglect the effect of the stabilization time of the conversion factor on the total composition 311 of the emissions and treat each cell of the grid independently from its neighbours. Beirle et al., 2019 [13] found an 312 annual average of 1.32 for this conversion factor, but CAMS data shows small deviations from this value over Egyptian 313 urban areas, as urban concentrations depend on local conditions. We therefore calculate for each cell of the domain 314 the following equation: 315

$$e_{\rm NO_x} = \frac{[\rm NO_x]}{[\rm NO_2]} e_{\rm NO_2} \tag{6}$$

For convenience, quantities  $\frac{[NO_x]}{[NO_2]} \operatorname{div}(\Omega_{NO_2}\mathbf{w})$  and  $\frac{[NO_x]}{[NO_2]}\Omega_{NO_2}/\tau$  represent the respective contributions of the transport and the sink terms to total NO<sub>x</sub> emissions. We obtain the emissions related to human activity at the scale of the country by removing the influence of the NO<sub>2</sub> background detected by TROPOMI and the possible non-anthropogenic sources of NO<sub>x</sub>. Neglecting the part of the country that lies outside the domain, the average surface emissions from Egypt's anthropogenic activity  $E_{NO_x}$  can be calculated as being the difference between NO<sub>x</sub> emissions over all urban cells and NO<sub>x</sub> emissions over all rural cells:

$$E_{\rm NO_x} = \frac{1}{n_{urb}} \sum_{i=1}^{n_{urb}} e_{\rm NO_x, i} - \frac{1}{n_{rur}} \sum_{j=1}^{n_{rur}} e_{\rm NO_x, j}$$
(7)

For robust statistics, these derived emissions can be averaged monthly, enabling a month-by-month comparison with 322 bottom-up inventories. The linearity of the averaging processes ensures the interchangeability of temporal and spatial 323 averages. A monthly average is relevant because it aggregates enough data to limit the high inter-day variability 324 due to changing wind patterns or differences between week days and week-ends. In addition, it enables the study 325 of monthly  $NO_x$  emission profiles which reflect changes in human activities throughout the year due to temperature 326 changes, economic constraints and cultural norms. The use of the average value for rural cells as a "background" 327 that is subtracted from the emission estimates for urban cells assumes that all natural processes in  $NO_x$  emissions 32 lead to similar amounts for both types of cells, which is not trivial, as the vast majority of rural cells are located 329 in desert, mountainous or oceanic areas while the vast majority of urban cells are located near the Nile valley or 330 the Mediterranean coast. However, as the processes involved in natural  $NO_x$  sources lead to emissions smaller than 331 anthropogenic emissions in polluted areas (e.g. Lin, 2012 [57] for China), we neglect this difference in the following 332 calculations. Finally, the total anthropogenic NO<sub>x</sub> emissions of the domain can be obtained by multiplying  $E_{\rm NO_x}$  by 333 the cumulative area covered by the urban cells (approximately  $95,000 \text{ km}^2$ ). 334

#### 335 4 Results and discussion

#### 336 4.1 Line densities and NO<sub>2</sub> lifetime

Here we compare the results of the TROPOMI line densities fits for Rivadh to the lifetime calculated by Equation 337 (2) using CAMS data. The two years of TROPOMI observations (from November 2018 to November 2020) provide a 338 wide variety of profiles. For level  $\mathcal{B}$ , Figure 4 also provides the number of samples that are being kept at each stage of 339 the process. Of the 731 observations available, 203 have a wind direction in the cone with a north-south orientation 340 with an aperture of  $40^{\circ}$  (i.e. between  $340^{\circ}$  and  $20^{\circ}$  or between  $160^{\circ}$  and  $200^{\circ}$ ). Of the remaining observations, 166 341 have a sufficiently clear sky to be retained. The criterion of weak variability for the wind direction brings to 91 the 342 number of observations that are kept by the method. On these 91 observations, the line density profiles are calculated 343 and the fits applied. According to Equation (5), the lifetime is calculated using the mean wind module around the 344 point source. The two lifetimes are calculated with the parameters taken at the medium level  $\mathcal{A}$  or at the top level 345 B. Of the 91 fits obtained, 52 are of high quality (correlation coefficient between fit function and line density profile 346 347 greater than 0.97) for level  $\mathcal{A}$  and 51 for level  $\mathcal{B}$ . 39 of these fits lead to a lifetime  $\tau_{\text{fit}}$  greater than 1.0 h for both levels (they do not necessarily correspond to the same days). All remaining samples correspond to atmospheric conditions 348 with moderate to fast winds, with a module ranging between 2 and 11 m/s (with an average of 5.9 m/s) for level A349 and between 3 and 8 m/s (with an average of 5.4 m/s) for level  $\mathcal{B}$ . 350







Figure 5: (left) Comparison between CAMS-induced NO<sub>2</sub> lifetimes and lifetimes from NO<sub>2</sub> line density fittings with EMG function above Riyadh city centre, for level  $\mathcal{A}$  (a) and  $\mathcal{B}$  (c). The samples presented correspond to patterns in clear-sky conditions for which the mean wind is in the north-south direction with a low variance, and for which the correlation between line density profile and fit gives a correlation coefficient of more than 0.97 and a lifetime of more than 1.0 h. No NO<sub>2</sub> patterns during the summers of 2019 and 2020 meet these conditions. Dashed lines separate the groups of observations by season. (right) Comparison between the two calculated lifetimes for level  $\mathcal{A}$  (b) and  $\mathcal{B}$  (d). A linear regression with an intercept forced to be zero is displayed with a green dashed line.

These lifetimes are compared to the corresponding lifetimes obtained from CAMS data in Figure 5 for each sample, 351 which is divided into seasons for a more convenient comparison. The use of level  $\mathcal{A}$  leads to a notable underestimation 352 of the NO<sub>2</sub> lifetime in autumn compared to the lifetime calculated with the fitting method. This same level leads to 353 an overestimation of the lifetime in winter. This trend is not found with the use of level  $\mathcal{B}$ , which leads to a better 354 reproduction of the lifetimes calculated with the fitting method for the available seasons. Figure 5 shows a linear 355 regression between the two calculated lifetimes. The results are scattered, with a correlation coefficient higher for 356 level  $\mathcal{B}$  (0.376) than for level  $\mathcal{A}$  (0.139). When the intercept of the regression line is forced to zero, the resulting slope 357 is closer to 1 for level  $\mathcal{B}$  (0.953) than for level  $\mathcal{A}$  (1.086). Level  $\mathcal{B}$  is therefore the one that leads to the best match 358 between the lifetime calculated with Equation (2) and the lifetime calculated from line densities. The results that 359 are presented in the following sections (except for Section 4.3) are therefore results of calculations performed with 360 parameters (w, [OH], T and  $[NO_x]/[NO_2]$ ) estimated at level  $\mathcal{B}$ . Nevertheless, it should be noted that no summer 361 observations were included in the comparison. The main reason for this is the wind direction: of the 188 summer 362 days observed, 178 of them have a mean wind direction outside the north-south cone over central Riyadh. On the 363 remaining ten days (one for summer 2019 and nine for summer 2020), the ERA5 wind direction is too variable for the 364 fit to be considered relevant, or the fit results in a correlation coefficient below 0.97. Thus it is not clear how correctly 365 the  $NO_2$  lifetime would be calculated during both summer periods by Equation (2). OH concentrations being the 366 main driver of this lifetime, we cannot assess the relevance of the representation of OH concentrations by CAMS data 367 during summer days in the study. 36

### <sup>369</sup> 4.2 Mapping of Egypt's NO<sub>x</sub> emissions

The top-down emission model is then applied to the Egyptian domain with CAMS OH concentration and temperature fields for lifetime calculations. For each cell,  $NO_x$  emissions are calculated according to Equation (6), resulting in a mapping of Egypt's emissions. The obtained values are averaged monthly from November 2018 to November 2020.





Figure 6 shows a composition of the emissions map with the transport term and the sink term for the months of 373 January and July 2019. The Nile appears on transport term maps: the divergence calculation complies with what is 374 expected from a line density of emitters, i.e. a clear separation of zones of positive divergence from zones of negative 375 divergence with a separation line corresponding to the course of the river. The fact that the negative divergence 376 377 zones are located to the east of the river and the positive divergence zones to the west is the result of the wind being predominantly towards in the northeast and southeast quadrants during the overflight of the region. Some 378 point sources like Cairo, Alexandria, Asyut or Aswan are easily identifiable. The sink term, directly proportional to 379 the TROPOMI slant column densities, also highlights these cities. However, unlike the transport term, which has 380 a similar spatial pattern from month to month, the sink term is clearly stronger in summer than in winter. This is 381 mainly due to a higher lifetime in winter than in summer (4.00 h on average in January 2019 and 2.15 h in July 2019) 382 while the average TROPOMI NO<sub>2</sub> concentrations are slightly higher during winter  $(1.687 \times 10^{15} \text{ molecules.cm}^{-2} \text{ for})$ 383 January 2019 and 1.440  $\times 10^{15}$  molecules.cm<sup>-2</sup> for July 2019). Over the whole domain, the mean transport term varies throughout the studied period between  $-0.035 \times 10^{15}$  molecules.cm<sup>-2</sup>.h<sup>-1</sup> (January 2019) and  $0.026 \times 10^{15}$ 384 385 molecules.cm<sup>-2</sup>.h<sup>-1</sup> (May 2020). Thus, it hardly contributes to the NO<sub>x</sub> emission budget, the mean chemical sink 386 term alone varying between  $0.451 \times 10^{15}$  molecules.cm<sup>-2</sup>.h<sup>-1</sup> (January 2020) and  $1.121 \times 10^{15}$  molecules.cm<sup>-2</sup>.h<sup>-1</sup> 387 (September 2020). 388



Figure 6: Construction of  $NO_x$  emissions above most of Egypt's territory: (top) transport term (A), sink term (B) and the resulting surface emissions, being counted as total emissions (C) for January 2019. (bottom) transport term (D), sink term (E) and the resulting surface emissions (F) for July 2019.

Several cities in the country are thus highlighted as the main emitters of the country, such as Cairo, Alexandria, Beni
Suef, Asyut or Aswan. The industrial area of Ain-Sokhna, located southwest of Suez, also appears as a main emitter.
Table 1 compares the monthly values for the sink term and the absolute value of the transport term above five major
cities of the country, with populations ranging from 193,000 to 19 million inhabitants, as well as Ain-Sokhna's area.
The mean values for TROPOMI column densities are also provided. According to the results, the capital city of Cairo





is by far the largest emitter in the country, largely due to its large population, resulting in high traffic emissions, but 394 also to its intensive industrial activity. Alexandria, the country's second largest city, is not necessarily the second 395 largest emitter, as its emissions are comparable to those of smaller cities such as Beni Suef or Asyut. However, the 396 three cities concentrate a large amount of industrial activity: Alexandria hosts several oil and gas power plants and a 397 398 small number of cement factories, while Beni Suef is close to several oil and gas power plants and hosts several flaring sites. Similarly, the city centre of Asyut is close to two oil and gas-fired power plants and a cement factory. This seems 399 to indicate that industrial activity might be the main factor of  $NO_x$  emissions differences between these cities, before 400 population size. This explains why  $NO_x$  emissions from these three cities are comparable to those of the industrial 401 area of Ain Sokhna, which includes several cement plants, iron smelters and oil and gas plants. It might also explain 402 why Aswan, which has a population that is comparable to Beni Suef or Asyut, but which does not have any major 403 industrial site, has slightly lower emissions than the two other cities. Finally, the Gulf of Suez displays relatively large 404 emissions, which might be attributed to the shipping sector, the region being a major gateway for international trade. 405 Because it also hosts several flaring sites, these emissions might also be due to the oil and gas extraction activity. 406

City	Population density	Ja	n. 2019		Jul. 2019		
	1 optiation density	$\Omega_{\rm NO_2}$	Transport	Sink	$\Omega_{ m NO_2}$	Transport	Sink
	$(\rm khab/km^2)$	$(\mathcal{M}_{\rm NO_2}.\rm cm^{-2})$	$(\mathcal{M}_{\rm NO_{\bf x}}.\rm cm^{-2}.h^{-1})$		$(\mathcal{M}_{\mathrm{NO}_2}.\mathrm{cm}^{-2})$	$(\mathcal{M}_{NO_x}.cm^{-2}.h^{-1})$	
Cairo	52.2	18.016	5.615	5.520	8.331	2.901	5.883
Alexandria	3.2	5.569	2.047	1.188	2.518	0.694	2.402
Asyut	3.0	4.134	1.230	1.298	4.358	2.041	3.110
Aswan	1.6	2.615	0.431	1.521	2.175	0.555	1.532
Beni Suef	2.5	7.472	0.974	2.513	4.683	1.571	4.238
Ain Sokhna	- (industrial area)	8.159	1.585	2.548	5.216	2.012	4.737

Table 1: Comparison between the transport term and the sink term above different cities among the 20 most populous cities in Egypt, as well as the industrial region of Ain Sokhna located 45 km southwest of Suez for January and July 2019. Numbers correspond to average values within 18 km from the city centre. The value for the mean TROPOMI NO<sub>2</sub> column density is also given. The population density of the corresponding governorate (2020 value) is noted as a comparison. Unit  $\mathcal{M}$  stands for a quantity of  $10^{15}$  molecules (NO<sub>2</sub> or NO<sub>x</sub>).

Although these cities and areas can be described as high-emission sites, the term responsible for these emissions differ 407 408 from one site to the other. Figure 7 shows the contribution of the transport term (taken in absolute value) to total emissions for January and July 2019. Because wind fields are relatively homogeneous along the Nile on spatial scales 409 of less than 100 km,  $NO_2$  concentration gradients perceived by TROPOMI in the region mainly contribute to the 410 increase of the transport term which can reach similar values as the sink term. Conversely, desert areas such as the 411 Libyan Desert, the Eastern Desert and the Sinai region, (located respectively to the west, east and northeast of the 412 Nile) show a very low value for the transport term compared to the sink term, due to the homogeneity of both the 413 wind field and the detected  $NO_2$  concentrations in these areas. Finally, a strong predominance of the transport term 414 can be observed near coastal regions even without the presence of emitters nearby due to changing winds. 415



Figure 7: Share of the absolute value of the transport term in the sum of the sink term and the absolute value of the transport term above most of Egypt's territory, indicating the local importance of the transport term in  $NO_x$  emissions above urban areas. The average of this ratio is shown for January 2019 (A) and July 2019 (B).





As a consequence, most cities in the country are characterised by a transport term which can locally be of the same 416 order of magnitude as the sink term, especially during the winter season. In the case of the Gulf of Suez, the transport 417 term can be 1 to 2 times higher than the sink term, which varies between 0.5 and  $1.0 \times 10^{15}$  molecules.cm<sup>-2</sup>.h<sup>-1</sup>. Those 418 values are slightly higher than the average emissions above rural cells areas due to sink term (about  $0.4 - 0.6 \times 10^{15}$ 419 molecules.cm<sup>-2</sup>.h<sup>-1</sup> in winter and  $0.7 - 0.9 \times 10^{15}$  molecules.cm<sup>-2</sup>.h<sup>-1</sup> in summer). It is also observed that TROPOMI 420 NO<sub>2</sub> column densities above this zone are relatively homogeneous, which indicates that the high value for the transport 421 term is due to abrupt changes in wind direction, which is consistent with the presence of the coast. Consequently, a 422 poor representation of the wind field by ERA5, caused by the  $0.25^{\circ} \times 0.25^{\circ}$  spatial resolution of the data (i.e. about 423 26 km in this region, the same order of magnitude as the width of the channel) might misestimate the transport term. 424

#### 425 4.3 Vertical analysis

Here we investigate the influence of the choice of the vertical level in the representation of the model parameters. 426 This influence is of considerable importance, as  $NO_x$  sources in urban areas can be located at different altitudes. For 427 instance, emissions from the road sector from tailpipes are located at ground level, whereas  $NO_x$  from power plants 428 and industrial facilities can be emitted from stacks, which are usually located between 50 and 300 m above the ground. 429 The results of Section 4.1 showed that level  $\mathcal{B}$  was more appropriate for the representation of the NO<sub>2</sub> lifetime. This 430 level is therefore chosen as a reference for the comparison. We study the effect of a transition from level  $\mathcal B$  to level 431  $\mathcal{A}$  for each of the 3 parameters involved in the representation of the sink term, i.e. temperature T, hydroxyl radical 432 concentration [OH] and concentration ratio  $[NO_x]/[NO_2]$ . The results for the averages over urban and rural areas are 433 given for the months of January, April, July and October 2019 in Table 2. As the wind field is only involved in the 434 transport term whose spatial integration nearly leads to zero, the influence of this parameter is not studied. 435

TITLE		$\mathbf{Sink \ term} \ (10^{15} \ \mathbf{molecules.cm}^{-2}.\mathbf{h}^{-1})$								
level $\mathcal{B}$	level $\mathcal{A}$	Jan. 19	Jan. 19	Apr. 19	Apr. 19	Jul. 19	Jul. 19	Oct. 19	Oct. 19	
(987.5 hPa)	(925 hPa)	(urban)	(rural)	(urban)	(rural)	(urban)	(rural)	(urban)	(rural)	
$T, [OH], \frac{[NO_x]}{[NO_2]}$	-	1.289	0.493	1.737	0.694	1.969	0.788	1.627	0.672	
IOHI [NO <sub>x</sub> ]	Т	1.349	0.513	1.827	0.727	2.073	0.826	1.709	0.704	
$[OII], \overline{[NO_2]}$	1	(+4.7%)	(+4.2%)	(+5.2%)	(+4.8%)	(+5.3%)	(+4.9%)	(+5.0%)	(+4.7%)	
$T, \frac{[NO_x]}{[NO_2]}$	[OH]	0.968	0.366	1.504	0.603	1.837	0.750	1.403	0.573	
		(-24.9%)	(-25.7%)	(-13.4%)	(-13.1%)	(-6.7%)	(-4.8%)	(-13.8%)	(-14.7%)	
T, [OH]	[NO <sub>x</sub> ]	1.310	0.499	1.776	0.712	1.999	0.803	1.659	0.693	
	$\overline{[NO_2]}$	(+1.7%)	(+1.4%)	(+2.2%)	(+2.5%)	(+1.5%)	(+1.9%)	(+2.0%)	(+3.1%)	

Table 2: Analysis of the effect of a vertical change in the parameters used to estimate the sink term in  $NO_x$  emissions: temperature, hydroxyl radical concentration, and  $NO_x:NO_2$  concentration ratio. The comparison is conducted between the estimated quantities for level  $\mathcal{B}$  and level  $\mathcal{A}$ . The comparison is conducted for four months of the year 2019. Values within brackets represent the variation from the base case for which all quantities are estimated at level  $\mathcal{B}$ .

The transition to the level  $\mathcal{A}$  generally results in a decrease in temperature, leading to an increase in the reaction rate 436  $k_{mean}$  and thus an increase in the emissions from the sink term. This transition has only a small influence on the 437 total  $NO_x$  emission estimates, with the total increasing by 4 to 6%. The effect is slightly more pronounced in urban 438 areas, due to a steeper vertical temperature profile in these areas. The influence of OH goes in the opposite direction: 439 its concentration decreases strongly with altitude, weakening the sink term. The share of emissions due to the sink 440 term being proportional to this parameter, the effect of the vertical is very pronounced. Thus, the transition to level 441  $\mathcal{A}$  weakens the sink term by 4 to 9% in summer (with an average of -6.03% for the months June/July/August) and 442 by 9 to 26% in winter (with an average of -15.70% for the months December/January/February). This weakening 443 seems more pronounced over urban areas than over rural areas from March to October, and more pronounced over 444 rural areas than over urban areas during the rest of the year. Finally, the influence of the  $NO_x:NO_2$  ratio is negligible 445 on the NO<sub>x</sub> emission estimates. Thus, the transition to level  $\mathcal{A}$  results in an increase in the sink term of 2 to 4%, 446 due to a decrease in both concentrations of NO and  $NO_2$  with respect to the vertical but with a greater decrease for 447  $NO_2$ . This vertical study confirms the crucial importance of the OH concentration representation for the accurate 448 representation of  $NO_x$  emissions. OH concentration appears here as the crucial driver of the sink term, much more 449 sensitive to vertical differences than temperature or the NO<sub>x</sub>:NO<sub>2</sub> concentration ratio. 450

#### 451 4.4 Weekly cycle

452 In Egypt, the official rest day is Friday, and the economic activity of the country is a priori lower during this day than





during the other days of the week. We therefore try to characterise this feature, by evaluating the weekly cycle of 453  $NO_x$  emissions. We use the TROPOMI-inferred emissions to obtain averages per day of the week. We use the quality 454 assurance  $q_a$  of TROPOMI retrievals to ignore the days for which more than 20% of the domain has low-quality 455 data (this happens 43 times in 2018/2019 and 28 times in 2019/2020). Figure 8 shows the resulting emissions for 456 457 the period November 2018 - November 2019 and November 2019 - November 2020. A Friday minimum is observed, defining a marked weekly cycle. This trend is also observed for mean  $NO_2$  column densities, for which no intra-weekly 458 variation is observed. Over the 2018-2019 period, Fridays have average emissions of  $0.978 \pm 0.408$  kt, which is lower than average emissions for the rest of the week, which reach  $1.279 \pm 0.533$  kt. A similar trend is observed in 2019-2020, 460 for which the average for Fridays is  $0.856 \pm 0.357$  kt and the average for other days is  $1.067 \pm 0.449$  kt. The difference 461 in emissions between the two periods is due to smaller emissions in December 2019, January 2020 and February 2020 462 that are discussed in Section 4.5. On average, Friday emissions correspond to a ratio of 0.81:7 (i.e. a value of 0.81 463 after normalisation on the seven days of the week) for the entire domain. This result is consistent with the values 464 obtained by Stavrakou et al., 2020 [58], who used TROPOMI data and another emission model to calculate a ratio of 465 0.71:7 for Cairo and 0.89:7 for Alexandria in 2017. 466



Figure 8: Weekly profiles of anthropogenic NO<sub>x</sub> emissions for Egypt using TROPOMI observations in 2018-2019 (purple line) and 2019-2020 (green line). Thin grey lines represent individual weeks. Days for which less than 80% of the domain counts low quality observations  $(q_a < 0.75)$  are not represented.

#### 467 4.5 Impacts of lockdown during COVID-19

The ongoing global outbreak of COVID-19 forced many countries around the world to implement unprecedented 468 public health responses, including travel restrictions, quarantines, curfews and lockdowns. Such measures have helped 469 to counter the spread of the virus and have, meanwhile, caused high reductions in global demand for fossil fuels (IEA, 470 2020 [59]). They also led to a fall in the levels of NO<sub>2</sub> and other air pollutants across the globe (Venter et al., 2020 471 [60]; Bauwens et al., 2020 [61]; Barré et al., 2021 [62]). To prevent the spread of COVID-19, Egyptian authorities 472 ordered a partial lockdown from March 15th till June 30th 2020, closing all public areas (e.g. sport centres, nightclubs, 473 restaurants and cafes) and suspending religious activities in all mosques and churches throughout the country. They 474 also implemented more drastic measures such as a full lockdown during Easter (April 20th) and Eid (May 23rd to 475 May 25th), resulting in before lifting some restrictions on June 1st (Hale et al., 2021 [63]). In addition to the effect 476 of containment on the activity of the country, the global decline in consumption led to a drop in the production of 477 certain industrial products. 478

Several studies have estimated the impact of these events on the air pollution levels in the urban centres of the country : from in-situ measurements, El-Sheekh et al., 2021 [64] estimated that NO<sub>2</sub> concentrations had dropped by 25.9% in Alexandria's city centre after the start of the lockdown on March 13th, while El-Magd et al., 2020 [65] used OMI retrievals to estimate a 45.5% reduction of NO<sub>2</sub> concentrations for the entire country during the spring compared to 2018 and 2019 average values. However, due to a changing lifetime of NO<sub>2</sub>, reductions in the concentrations of NO<sub>2</sub> might not be entirely due to a drop of NO<sub>x</sub> emissions, which leads us to focus on the variation of NO<sub>x</sub> emissions during this singular period. Using our top-down emission model, reductions in total NO<sub>x</sub> emissions of 34.6%, 17.4%





and 16.6% are observed for the respective months of March, April and May 2020 compared to the equivalent months 486 in 2019. This drop of emissions in 2020 compared to 2019 calculated by the model also correspond to a decrease in 487 observed NO<sub>2</sub> columns. However, no significant changes in OH concentrations seem to appear: on average, from 2019 488 to 2020, CAMS near-real-time data shows a decrease of 5.5% for OH concentration over the urban cells for the period 489 March/April/May, while TROPOMI retrievals above urban areas show a decrease in NO<sub>2</sub> column densities of 21.6% 490 over the same period. However, these effects observed for the months of March, April and May 2020 are not repeated 491 in June 2020, for which emissions show an increase of 12.3% compared to June 2019. This increase is largely the 492 result of an increase in the difference between urban and rural average emissions (as calculated according to Equation 493 7). Indeed, the urban term of June 2020 is higher than that of June 2019, due to an increase in TROPOMI urban 494 concentrations (+2.4%) while the NO<sub>2</sub> lifetime is almost unchanged (+0.4%). The rural term varies in the opposite 495 direction: a decrease in TROPOMI rural concentrations (-7.6%) is observed while NO<sub>2</sub> lifetime increases strongly 496 (+11.5%). This increase in June emissions seems to indicate that the lift on restrictions allowed a catch-up of the 497 economic activity which has be sufficiently strong to generate higher emissions in 2020 than in 2019. 498

#### 499 4.6 Annual cycle and comparison to inventories

Here, we attempt to compare our TROPOMI-derived NO<sub>x</sub> emissions to emissions from CAMS and EDGAR inventories. 500 Figure 9 shows the total anthropogenic  $NO_{x}$  emissions over the urban cells from November 2018 to November 2020 (i.e. 501 a period of two years), with the average anthropogenic emissions calculated according to Equation (7). As indicated in 502 Section 3.2, the emissions, calculated at 13:30 local time, are representative of the average daily consumption in Egypt. 503 The total calculated for each month therefore corresponds to the  $NO_x$  production by human activities in Egyptian 504 urban and industrial areas. After aggregating the different sectors of activity, CAMS and EDGAR inventories directly 505 provide the anthropogenic  $NO_x$  emissions over the same areas. All  $NO_x$  emissions are expressed in mass terms as NO. 506 We note that the EDGAR inventory does not cover the period 2018-2020 (the inventory ends in 2015). On Figure 9, 507 EDGAR emissions corresponding to the period between November 2013 and November 2015 are displayed, i.e. with 508 a delay of 5 years compared to TROPOMI-derived emissions and CAMS estimates. TROPOMI-derived emissions 509 are lower than the inventory estimates between November 2019 and February 2020. In particular, the difference is 510 significant in December 2019 and January 2020 (respectively 54.8% and 55.7% of CAMS levels). In the computations, 511 this decrease is mainly due to a relatively low value of the OH concentration which reaches  $5.86 \times 10^6$  molecules.cm<sup>-3</sup> 512 on average for these two months, with  $4.95 \times 10^6$  molecules.cm<sup>-3</sup> above urban areas and  $6.09 \times 10^6$  molecules.cm<sup>-3</sup> 513 over rural areas. They were respectively 6.96, 6.94 and 6.97  $\times 10^6$  molecules.cm<sup>-3</sup> for the previous year (December 514 2018-January 2019) and 7.24, 6.94 and 7.31  $\times 10^6$  molecules.cm<sup>-3</sup> for the year before (December 2017-January 2018). 515 A decrease in tropospheric columns (-14.3% for urban areas and -4.6% for rural areas) also contributes to this drop. 516



Figure 9: Comparison of TROPOMI-derived  $NO_x$  emissions for urban areas in Egypt (light blue), with the corresponding emissions from EDGAR (red with stripes) and CAMS (yellow) inventories. EDGAR data is provided for comparison purposes and covers the years 2013-2015. Error bars for TROPOMI-derived emissions are calculated using uncertainties for the parameters involved in Equation (3).





Except for this singular period, TROPOMI-derived indications of emissions are higher than the CAMS inventory 517 estimates. The top-down model estimates total emissions of 814.5 kt over the 24 months, which is 162.9 kt higher 518 than CAMS for the same period (651.6 kt). Of these 162.9 kt, 125.2 kt are emitted during the first 12 months 519 (before the underestimation period) and 37.7 kt during the 12 following months. The average value for top-down 520  $NO_x$  emissions are 25.0% higher than CAMS estimates. TROPOMI-inferred emissions show an annual variability: 521 the emissions seem to follow a one-year seasonal cycle where emissions are higher in summer than in winter. These 522 results, at first sight, seem to be correlated with power emissions which dominate the use of fossil fuels in Egypt 523 (Abdallah et al., 2020 [66]). These power emissions are due to the country's residential electricity consumption (Attia 524 et al., 2012 [49]; Elharidi et al., 2013 [67]; Nassief, 2014 [68]). They also meet the needs of industry. Summer peaks in 525 electricity consumption are mostly driven by temperature: for instance, the sales of air conditioning and ventilation 526 systems have been increasing for several decades (Wahba et al., 2018 [69]). The use of air conditioning in cars, which 527 requires an additional amount of fuel, could also contribute to the increase of  $NO_x$  emissions in summer. However, 528 some features of the industrial activities in the region might be counteracting this trend. For some sectors such as 529 cement or steel, production is lower in summer, due to the physical wear experienced by workers due to heat, but 530 also due to certain periods of leave. Given the importance of industrial activities in the production of  $NO_x$  shown in 531 Section 4.2, this aspect cannot be neglected. The transport sector could also counteract the observed trend: although 532 the use of air conditioning in cars increase  $NO_x$  emissions of the sector, the observed mean traffic in the country is 533 higher between November and February and lower between June and August, especially in Cairo which gathers most 534 of the population. In the absence of additional data, it therefore seems difficult to conclude on the relevance of the 53 seasonal cycle that seems to be produced by our top-down model. This caution is all the more necessary as CAMS 53( and EDGAR inventories show different seasonal cycles in  $NO_x$  emissions, with different dynamics: while the EDGAR 537 inventory predicts a maximum of emissions in December or January and a minimum in April, the CAMS inventory 538 shows two local maxima each year in May and November and two local minima in February and September. The amplitude of the cycle is higher in EDGAR than in CAMS, representing 14.2% of the average value for emissions 540 estimates in EDGAR and 12.4% in CAMS. These differences between the model and the inventories do not give us 541 any information on the seasonality of  $NO_x$  emissions that should be found in the outputs of our top-down model. 542

#### 543 4.7 Uncertainties and assessments of the previous results

The estimation of  $NO_x$  emissions is based on the use of several quantities with varying uncertainties. The error bars 544 shown in Figures 5 and 9 are thus calculated from uncertainty statistics whose references are presented in this section. 545 Since these references do not specify the exact nature of these statistics, we assume they correspond to standard 546 deviations. The uncertainty of tropospheric  $NO_2$  columns under polluted conditions is dominated by the sensitivity of satellite observations to lower tropospheric air masses, expressed by the tropospheric air-mass factor (AMF). The AMF depends on the viewing geometry, surface albedo, NO<sub>2</sub> vertical profile, and cloud characteristics (Lorente et al., 2017 [38]; Eskes et al., 2019 [18]). The column relative uncertainty due to the AMF is of the order of 30% (Boersma et 550 al., 2004 [37]). S-5P validation activities indicate that TROPOMI tropospheric NO<sub>2</sub> columns are systematically biased 551 low by about 30%-50% over cities, which is most likely related to the *a priori* profiles used within the operational 552 retrieval that do not reflect well the NO<sub>2</sub> peak close to ground. For the Middle East region, the impact of the *a priori* 553 profile is less critical, as surface albedo is generally high and cloud fractions are generally low. Thus, we expect no 554 such bias, and consider a relative uncertainty of 30% for the tropospheric column. Other uncertainties must be taken 555 into account: the transition from  $NO_2$  TROPOMI columns to  $NO_x$  emissions requires parameters which appear in 556 Equation (2) and Equation (3). For both zonal and meridional wind components, we assume an uncertainty of 1 m/s557 (Coburn et al., 2019 [70]). For [OH], the analysis of different methods conducted by Huijnen et al., 2019 [71] showed 558 smaller differences for low latitudes than for extratropics, but still significant. We thus take a relative uncertainty of 559 560 30% for OH concentration. For the reaction rate  $k_{mean}$ , the value of the corresponding relative uncertainty has been estimated by Burkholder et al., 2020 [28]. Because the sensitivity test conducted in Section 4.3 shows that changing 561 the temperature vertically only changes the results by 2-3%, and because vertical temperature gradients are much 562 stronger than horizontal temperature gradients, then the uncertainty related to the horizontal temperature field is 563 small. Therefore, we neglect the impact of temperature on final uncertainty. As a consequence, the propagation of 564 these different uncertainties on the monthly estimates of  $NO_x$  emissions in Egypt leads to an expanded uncertainty 565 between 40 and 43%. For lifetimes calculated with the EMG function fitting, the corresponding expanded uncertainty 566 ranges between 18% and 79%. 567

We acknowledge the fact that our treatment of uncertainties is simplified there. Many minor sinks highlighted in Section 3.1 are neglected in the calculations, and the corresponding uncertainty are not taken into account. Moreover, among the remaining sources of uncertainties, the major ones are treated with fixed values for relative uncertainties,





which leads to absolute uncertainties that are roughly proportional to monthly emissions. As a result, the confidence interval displayed on Figure 9 is larger in summer than in winter (with a length of 6.0 kt in January 2020 and of 17.4 kt for July 2020), and the drop in emissions for winter 2019-2020 appears as a persistent feature of the model outputs. If this drop is realistic, then our top-down model provides a method for improving national inventories. If it is not, then several assumptions of our model can be questioned. For instance, because this drop is largely due to variations in OH concentrations provided by CAMS, whose reliability has been evaluated for Riyadh, then the transposability hypothesis between Riyadh and Egypt may be subject to wider discussion. A better understanding of OH levels in Egypt, supported by in-situ measurements, might answer these questions and allow to improve our model.

#### 579 5 Conclusions

In this study, we investigated the potential of a top-down model of  $NO_x$  emissions based on TROPOMI retrievals at 580 high resolution over Egypt. The model is based on the study of a transport term and a sink term that requires different 581 parameters to be calculated. Among those parameters, the concentration in OH, involved in the calculation of the  $NO_2$ 582 mixed lifetime, is of fundamental importance. The comparison between the two ways of calculating the lifetime of  $NO_2$ 583 shows that OH concentration provided by CAMS data is reasonably reliable for the country. Parameters are therefore 584 taken in the first 200 m of the planetary boundary layer, because it is where OH shows the best consistency. However, the vertical sensitivity linked to this parameters remains high. Results illustrate the importance of the transport 586 term at local scale, which is the same order of magnitude as the sink term above large cities and industrial facilities; 587 it ceases to be relevant only at the country's scale. The top-down model is able to characterise declines in human 588 activities, whether they are due to restrictions during the COVID-19 pandemic or to Friday rest. It also estimates 589 higher emissions during summer. These high emissions might be interpreted by a higher consumption of electricity 590 driven by air-conditioning during hot days, but it remains unclear whether this pattern clearly reproduces changes in 591 human activity, in particular because the different emission inventories show different seasonalities. These inventories 592 also differ in the amount of total emissions: the average value for TROPOMI-derived  $NO_x$  emissions are 25.0% higher 593 than CAMS estimates. This discrepancy could be solved comparing the results of the model and inventory estimates to 594 industrial production or electricity consumption data at the scale of countries or regions. This study demonstrates the 595 596 potential of TROPOMI data for evaluating  $NO_x$  emissions in the EMME region. More generally, it demonstrates the importance of the contribution of independent observation systems to overcome the weaknesses of emission inventories, 597 provided that the local chemistry is well understood. The development of similar applications for different species is likely to allow better monitoring of global anthropogenic emissions, therefore helping companies and countries to 599 report their anthropogenic emissions of air pollutants and greenhouse gases as part of their strategies to tackle air 600 pollution issues and climate change. 601

#### 602 Data availability.

- $\mathbf{GAMS}\ \mathbf{NRT:}\ \mathbf{https://ads.atmosphere.copernicus.eu/cdsapp!/dataset/cams-global-atmospheric-composition-forecasts}$
- $\mathbf{ERA5}\ \mathbf{reanalysis:}\ \mathbf{https://cds.climate.copernicus.eu/cdsapp!/dataset/reanalysis-era5-pressure-levels-monthly-means}$
- 605 Global Rural-Urban Mapping Project (GRUMP): https://sedac.ciesin.columbia.edu/data/collection/grump-v1
- 606 Oil and gas power plants: http://globalenergyobservatory.org/
- 607 Industrial facilities: https://www.industryabout.com
- <sup>608</sup> Flaring sites: https://eogdata.mines.edu/download global flare.html
- 609 CAMS-GLOB-ANT v4.2: https://permalink.aeris-data.fr/CAMS-GLOB-ANT
- 610 EDGARv5.0: https://edgar.jrc.ec.europa.eu/dataset ap50
- **Competing interests.** The authors declare that they have no conflict of interest.
- Acknowledgements. The authors would like to thank Steven J. Davis (University of California, Irvine) and Dan
- Tong (Tsinghua University) for their contribution to the construction of our emitters database.
- **Financial support.** This study has been funded by the European Union's Horizon 2020 research and innovation programme under grant agreement N<sup>o</sup> 856612 (EMME-CARE).





## 616 References

- [1] A. Baklanov, L. T. Molina, and M. Gauss, "Megacities, air quality and climate," Atmospheric Environment, vol. 126, pp. 235–249, 2016.
- [2] A. Singh and M. Agrawal, "Acid rain and its ecological consequences," Journal of Environmental Biology, vol. 29, no. 1, p. 15, 2007.
- [3] U. EPA, "Integrated science assessment for oxides of nitrogen-health criteria," US Environmental Protection
   Agency, Washington, DC [Google Scholar], 2016.
- [4] M. Crippa, G. Janssens-Maenhout, F. Dentener, D. Guizzardi, K. Sindelarova, M. Muntean, R. Van Dingenen,
   and C. Granier, "Forty years of improvements in European air quality: regional policy-industry interactions with
   global impacts," Atmospheric Chemistry and Physics, vol. 16, no. 6, pp. 3825–3841, 2016.
- A. El-Magd, N. Zanaty, E. M. Ali, H. Irie, A. I. Abdelkader, et al., "Investigation of aerosol climatology, optical characteristics and variability over Egypt based on satellite observations and in-situ measurements," Atmosphere, vol. 11, no. 7, p. 714, 2020.
- 629 [6] UNEP (United Nations Environment Programme), "Air quality policies in Egypt," 2015.
- [7] B. Xue and W. Ren, "China's uncertain CO2 emissions," Nature Climate Change, vol. 2, no. 11, pp. 762–762, 2012.
- [8] C. Leue, M. Wenig, T. Wagner, O. Klimm, U. Platt, and B. Jähne, "Quantitative analysis of NOx emissions from global ozone Monitoring Experiment satellite image sequences," *Journal of Geophysical Research: Atmospheres*, vol. 106, no. D6, pp. 5493–5505, 2001.
- [9] R. V. Martin, D. J. Jacob, K. Chance, T. P. Kurosu, P. I. Palmer, and M. J. Evans, "Global inventory of nitrogen oxide emissions constrained by space-based observations of NO2 columns," *Journal of Geophysical Research: Atmospheres*, vol. 108, no. D17, 2003.
- [10] B. Mijling and R. Van Der A, "Using daily satellite observations to estimate emissions of short-lived air pollutants
   on a mesoscopic scale," *Journal of Geophysical Research: Atmospheres*, vol. 117, no. D17, 2012.
- [11] B. de Foy, Z. Lu, D. G. Streets, L. N. Lamsal, and B. N. Duncan, "Estimates of power plant NOx emissions and lifetimes from OMI NO2 satellite retrievals," *Atmospheric Environment*, vol. 116, pp. 1–11, 2015.
- [12] D. L. Goldberg, Z. Lu, D. G. Streets, B. de Foy, D. Griffin, C. A. McLinden, L. N. Lamsal, N. A. Krotkov, and
   H. Eskes, "Enhanced Capabilities of TROPOMI NO2: Estimating NOx from North American Cities and Power
   Plants," *Environmental science & technology*, vol. 53, no. 21, pp. 12594–12601, 2019.
- [13] S. Beirle, C. Borger, S. Dörner, A. Li, Z. Hu, F. Liu, Y. Wang, and T. Wagner, "Pinpointing nitrogen oxide emissions from space," *Science advances*, vol. 5, no. 11, p. eaax9800, 2019.
- [14] A. Lorente, K. Boersma, H. Eskes, J. Veefkind, J. Van Geffen, M. De Zeeuw, H. D. van der Gon, S. Beirle, and
   M. Krol, "Quantification of nitrogen oxides emissions from build-up of pollution over Paris with TROPOMI,"
   Scientific reports, vol. 9, no. 1, pp. 1–10, 2019.
- [15] K. Lange, A. Richter, and J. P. Burrows, "Variability of nitrogen oxide emission fluxes and lifetimes estimated
   from Sentinel-5P TROPOMI observations," *Atmospheric Chemistry and Physics Discussions*, pp. 1–32, 2021.
- [16] J. Veefkind, I. Aben, K. McMullan, H. Förster, J. De Vries, G. Otter, J. Claas, H. Eskes, J. De Haan, Q. Kleipool,
   *et al.*, "TROPOMI on the ESA Sentinel-5 Precursor: A GMES mission for global observations of the atmospheric composition for climate, air quality and ozone layer applications," *Remote sensing of environment*, vol. 120,
   pp. 70–83, 2012.
- [17] S. Lama, S. Houweling, K. F. Boersma, H. Eskes, I. Aben, H. A. Denier van der Gon, M. C. Krol, H. Dolman, T. Borsdorff, and A. Lorente, "Quantifying burning efficiency in megacities using the NO2/CO ratio from
   <sup>657</sup> The second sec
- the Tropospheric Monitoring Instrument (TROPOMI)," Atmospheric Chemistry and Physics, vol. 20, no. 17, pp. 10295–10310, 2020.





- [18] H. Eskes, K. Eichmann, J. Lambert, D. Loyola, J. Veefkind, A. Dehn, and C. Zehner, "S5P Mission Performance
   Centre Nitrogen Dioxide [l2\_no2\_] readme," Royal Netherlands Meteorological Institute (KNMI) De Bilt, the
   Netherlands, version, vol. 1, no. 00, 2019.
- [19] H. Hersbach, B. Bell, P. Berrisford, S. Hirahara, A. Horányi, J. Muñoz-Sabater, J. Nicolas, C. Peubey, R. Radu,
   D. Schepers, et al., "The ERA5 global reanalysis," Quarterly Journal of the Royal Meteorological Society, vol. 146,
   no. 730, pp. 1999–2049, 2020.
- [20] V. Huijnen, H. Eskes, A. Wagner, M. Schulz, Y. Christophe, M. Ramonet, S. Basart, A. Benedictow, A.-M.
   Blechschmidt, S. Chabrillat, *et al.*, "Validation report of the CAMS near-real-time global atmospheric composition service: System evolution and performance statistics. Status up to 1 June 2016.," 2016.
- 669 [21] J. H. Seinfeld, "Urban air pollution: state of the science," Science, vol. 243, no. 4892, pp. 745–752, 1989.
- [22] J. A. Logan, M. J. Prather, S. C. Wofsy, and M. B. McElroy, "Tropospheric chemistry: a global perspective," Journal of Geophysical Research: Oceans, vol. 86, no. C8, pp. 7210–7254, 1981.
- [23] H. Levy, "Normal atmosphere: Large radical and formaldehyde concentrations predicted," Science, vol. 173, no. 3992, pp. 141–143, 1971.
- [24] L. Valin, A. Russell, R. Hudman, and R. Cohen, "Effects of model resolution on the interpretation of satellite
   NO2 observations," *Atmospheric Chemistry and Physics*, vol. 11, no. 22, pp. 11647–11655, 2011.
- [25] J. Lelieveld, S. Gromov, A. Pozzer, and D. Taraborrelli, "Global tropospheric hydroxyl distribution, budget and
   reactivity," Atmospheric Chemistry and Physics, vol. 16, no. 19, pp. 12477–12493, 2016.
- [26] M. Li, E. Karu, C. Brenninkmeijer, H. Fischer, J. Lelieveld, and J. Williams, "Tropospheric OH and stratospheric OH and Cl concentrations determined from CH4, CH3Cl, and SF6 measurements," NPJ Climate and Atmospheric Science, vol. 1, no. 1, pp. 1–7, 2018.
- [27] G. M. Wolfe, J. M. Nicely, J. M. S. Clair, T. F. Hanisco, J. Liao, L. D. Oman, W. B. Brune, D. Miller, A. Thames,
  G. G. Abad, et al., "Mapping hydroxyl variability throughout the global remote troposphere via synthesis of
  airborne and satellite formaldehyde observations," *Proceedings of the National Academy of Sciences*, vol. 116,
  no. 23, pp. 11171–11180, 2019.
- [28] J. Burkholder, S. Sander, J. Abbatt, J. Barker, C. Cappa, J. Crounse, T. Dibble, R. Huie, C. Kolb, M. Kurylo,
   *et al.*, "Chemical kinetics and photochemical data for use in atmospheric studies; evaluation number 19," tech.
   rep., Pasadena, CA: Jet Propulsion Laboratory, National Aeronautics and Space ..., 2020.
- [29] K. Boersma, H. Eskes, E. Meijer, and H. Kelder, "Estimates of lightning NOx production from gome satellite observations," *Atmospheric Chemistry and Physics*, vol. 5, no. 9, pp. 2311–2331, 2005.
- [30] J. Yienger and H. Levy, "Empirical model of global soil-biogenic NOx emissions," Journal of Geophysical Research: Atmospheres, vol. 100, no. D6, pp. 11447–11464, 1995.
- [31] CIESIN, "CIESIN. Global Rural-Urban Mapping Project, Version 1 (GRUMPv1). Center for International Earth
   Science Information Network CIESIN Columbia University, International Food Policy Research Institute IFPRI, The World Bank, and Centro Internacional de Agricultura Tropical CIAT. 2011, Palisades, NY: NASA
   Socioeconomic Data and Applications Center (SEDAC)," 2019.
- [32] C. D. Elvidge, M. Zhizhin, K. Baugh, F.-C. Hsu, and T. Ghosh, "Methods for global survey of natural gas flaring from visible infrared imaging radiometer suite data," *Energies*, vol. 9, no. 1, p. 14, 2016.
- [33] M. Crippa, G. Oreggioni, D. Guizzardi, M. Muntean, E. Schaaf, E. Lo Vullo, E. Solazzo, F. Monforti-Ferrario,
   J. G. Olivier, and E. Vignati, "Fossil CO2 and GHG emissions of all world countries," *Publication Office of the European Union: Luxemburg*, 2019.
- 701 [34] C. Granier, S. Darras, H. D. van der Gon, D. Jana, N. Elguindi, G. Bo, G. Michael, G. Marc, J.-P. Jalkanen,
- J. Kuenen, et al., The Copernicus atmosphere monitoring service global and regional emissions (April 2019 version). PhD thesis, Copernicus Atmosphere Monitoring Service, 2019.





- [35] R. M. Hoesly, S. J. Smith, L. Feng, Z. Klimont, G. Janssens-Maenhout, T. Pitkanen, J. J. Seibert, L. Vu, R. J.
  Andres, R. M. Bolt, et al., "Historical (1750–2014) anthropogenic emissions of reactive gases and aerosols from the community emissions data system (CEDS)," *Geoscientific Model Development*, vol. 11, no. 1, pp. 369–408, 2018.
- [36] G. Janssens-Maenhout, M. Crippa, D. Guizzardi, M. Muntean, E. Schaaf, F. Dentener, P. Bergamaschi,
   V. Pagliari, J. G. Olivier, J. A. Peters, *et al.*, "Edgar v4. 3.2 global atlas of the three major greenhouse gas
   emissions for the period 1970–2012," *Earth System Science Data*, vol. 11, no. 3, pp. 959–1002, 2019.
- <sup>711</sup> [37] K. Boersma, H. Eskes, and E. Brinksma, "Error analysis for tropospheric NO2 retrieval from space," *Journal of Geophysical Research: Atmospheres*, vol. 109, no. D4, 2004.
- [38] A. Lorente, K. Folkert Boersma, H. Yu, S. Dörner, A. Hilboll, A. Richter, M. Liu, L. N. Lamsal, M. Barkley,
   I. D. Smedt, et al., "Structural uncertainty in air mass factor calculation for NO2 and HCHO satellite retrievals,"
   Atmospheric Measurement Techniques, vol. 10, no. 3, pp. 759–782, 2017.
- [39] S. Sander, R. Friedl, J. Abbatt, J. Barker, J. Burkholder, D. Golden, C. Kolb, M. Kurylo, G. Moortgat, P. Wine,
   *et al.*, "Chemical kinetics and photochemical data for use in atmospheric studies, jpl publication 10-6," *Evaluation*,
   no. 17, 2011.
- [40] E. R. Delaria, B. K. Place, A. X. Liu, and R. C. Cohen, "Laboratory measurements of stomatal NO2 deposition to native california trees and the role of forests in the NOx cycle," *Atmospheric Chemistry and Physics*, vol. 20, no. 22, pp. 14023–14041, 2020.
- [41] N. Sobanski, J. Thieser, J. Schuladen, C. Sauvage, W. Song, J. Williams, J. Lelieveld, and J. N. Crowley, "Day and night-time formation of organic nitrates at a forested mountain site in south-west germany," *Atmospheric Chemistry and Physics*, vol. 17, no. 6, pp. 4115–4130, 2017.
- [42] P. S. Romer Present, A. Zare, and R. C. Cohen, "The changing role of organic nitrates in the removal and transport of NOx," *Atmospheric Chemistry and Physics*, vol. 20, no. 1, pp. 267–279, 2020.
- [43] N. Butkovskaya, A. Kukui, N. Pouvesle, and G. Le Bras, "Formation of nitric acid in the gas-phase HO2+NO
   reaction: Effects of temperature and water vapor," *The Journal of Physical Chemistry A*, vol. 109, no. 29, pp. 6509–6520, 2005.
- [44] N. Butkovskaya, M.-T. Rayez, J.-C. Rayez, A. Kukui, and G. Le Bras, "Water vapor effect on the HNO3 yield in the HO2+NO reaction: experimental and theoretical evidence," *The Journal of Physical Chemistry A*, vol. 113, no. 42, pp. 11327–11342, 2009.
- [45] W. Moxim, H. Levy, and P. Kasibhatla, "Simulated global tropospheric PAN: Its transport and impact on NOx,"
   Journal of Geophysical Research: Atmospheres, vol. 101, no. D7, pp. 12621–12638, 1996.
- [46] C. A. Longfellow, A. Ravishankara, and D. R. Hanson, "Reactive uptake on hydrocarbon soot: Focus on NO2," Journal of Geophysical Research: Atmospheres, vol. 104, no. D11, pp. 13833–13840, 1999.
- [47] T. Stavrakou, J.-F. Müller, K. Boersma, R. Van Der A, J. Kurokawa, T. Ohara, and Q. Zhang, "Key chemical NOx sink uncertainties and how they influence top-down emissions of nitrogen oxides," *Atmospheric Chemistry and Physics*, vol. 13, no. 17, pp. 9057–9082, 2013.
- [48] N. Friedrich, P. Eger, J. Shenolikar, N. Sobanski, J. Schuladen, D. Dienhart, B. Hottmann, I. Tadic, H. Fischer,
   M. Martinez, et al., "Reactive nitrogen around the arabian peninsula and in the mediterranean sea during the
   2017 AQABA ship campaign," Atmospheric Chemistry and Physics, vol. 21, no. 10, pp. 7473–7498, 2021.
- [49] S. Attia, A. Evrard, and E. Gratia, "Development of benchmark models for the Egyptian residential buildings sector," *Applied Energy*, vol. 94, pp. 270–284, 2012.
- [50] EEHC, "Egyptian Electricity Holding Company annual report 2019/2020," 2021.
- [51] F. Rohrer and H. Berresheim, "Strong correlation between levels of tropospheric hydroxyl radicals and solar ultraviolet radiation," *Nature*, vol. 442, no. 7099, pp. 184–187, 2006.
- [52] S. Beirle, K. F. Boersma, U. Platt, M. G. Lawrence, and T. Wagner, "Megacity emissions and lifetimes of nitrogen oxides probed from space," *Science*, vol. 333, no. 6050, pp. 1737–1739, 2011.





- 750 [53] L. Valin, A. Russell, and R. Cohen, "Variations of OH radical in an urban plume inferred from NO2 column measurements," *Geophysical Research Letters*, vol. 40, no. 9, pp. 1856–1860, 2013.
- [54] M. Filioglou, E. Giannakaki, J. Backman, J. Kesti, A. Hirsikko, R. Engelmann, E. O'Connor, J. T. Leskinen,
   X. Shang, H. Korhonen, *et al.*, "Optical and geometrical aerosol particle properties over the United Arab Emirates,"
   *Atmospheric Chemistry and Physics*, vol. 20, no. 14, pp. 8909–8922, 2020.
- [55] T. Graedel, L. Farrow, and T. Weber, "Kinetic studies of the photochemistry of the urban troposphere," Atmo-
- <sup>756</sup> spheric Environment (1967), vol. 10, no. 12, pp. 1095–1116, 1976.
- 757 [56] J. H. Seinfeld and S. N. Pandis, "Atmospheric chemistry and physics from air pollution to climate change," 2006.
- [57] J.-T. Lin, "Satellite constraint for emissions of nitrogen oxides from anthropogenic, lightning and soil sources over
   East China on a high-resolution grid," *Atmospheric Chemistry and Physics*, vol. 12, no. 6, pp. 2881–2898, 2012.
- [58] T. Stavrakou, J.-F. Müller, M. Bauwens, K. Boersma, and J. van Geffen, "Satellite evidence for changes in the NO2 weekly cycle over large cities," *Scientific reports*, vol. 10, no. 1, pp. 1–9, 2020.
- [59] U. IEA, "Global energy review 2020," Ukraine. [Online] https://www. iea. org/countries/ukraine [Accessed: 2020-09-10], 2020.
- [60] Z. S. Venter, K. Aunan, S. Chowdhury, and J. Lelieveld, "COVID-19 lockdowns cause global air pollution declines,"
   *Proceedings of the National Academy of Sciences*, vol. 117, no. 32, pp. 18984–18990, 2020.
- [61] M. Bauwens, S. Compernolle, T. Stavrakou, J.-F. Müller, J. Van Gent, H. Eskes, P. F. Levelt, R. van der A,
   J. Veefkind, J. Vlietinck, et al., "Impact of coronavirus outbreak on NO2 pollution assessed using TROPOMI and
   OMI observations," Geophysical Research Letters, vol. 47, no. 11, p. e2020GL087978, 2020.
- [62] J. Barré, H. Petetin, A. Colette, M. Guevara, V.-H. Peuch, L. Rouil, R. Engelen, A. Inness, J. Flemming,
   C. Pérez García-Pando, *et al.*, "Estimating lockdown-induced european NO2 changes using satellite and surface observations and air quality models," *Atmospheric Chemistry and Physics*, vol. 21, no. 9, pp. 7373–7394, 2021.
- T. Hale, N. Angrist, R. Goldszmidt, B. Kira, A. Petherick, T. Phillips, S. Webster, E. Cameron-Blake, L. Hallas,
   S. Majumdar, et al., "A global panel database of pandemic policies (oxford covid-19 government response tracker),"
   Nature Human Behaviour, vol. 5, no. 4, pp. 529–538, 2021.
- [64] M. M. El-Sheekh and I. A. Hassan, "Lockdowns and reduction of economic activities during the COVID-19
   pandemic improved air quality in Alexandria, Egypt," *Environmental Monitoring and Assessment*, vol. 193, no. 1, pp. 1–7, 2021.
- [65] I. Abou El-Magd and N. Zanaty, "Impacts of short-term lockdown during COVID-19 on air quality in Egypt,"
   The Egyptian Journal of Remote Sensing and Space Science, 2020.
- [66] L. Abdallah and T. El-Shennawy, "Evaluation of CO2 emission from Egypt's future power plants," *Euro-*Mediterranean Journal for Environmental Integration, vol. 5, no. 3, pp. 1–8, 2020.
- [67] A. M. A. H. Elharidi, P. G. Tuohy, and M. Teamah, "Facing the growing problem of the electric power consumption in Egyptian residential building using building performance simulation program," in *Building simulation Cairo* 2013 conference, 2013.
- [68] M. M. Nassief, "Evaluation of electricity consumption of a residential flat in egypt," American Journal of Electrical Power and Energy Systems, vol. 3, no. 2, pp. 7–44, 2014.
- [69] S. M. Wahba, B. A. Kamel, K. M. Nassar, and A. S. Abdelsalam, "Effectiveness of green roofs and green walls on energy consumption and indoor comfort in arid climates," *Civil Engineering Journal*, vol. 4, no. 10, pp. 2284–2295, 2018.
- [70] J. J. Coburn, "Assessing wind data from reanalyses for the upper midwest," Journal of Applied Meteorology and Climatology, vol. 58, no. 3, pp. 429–446, 2019.
- [71] V. Huijnen, A. Pozzer, J. Arteta, G. Brasseur, I. Bouarar, S. Chabrillat, Y. Christophe, T. Doumbia, J. Flemming,
   J. Guth, et al., "Quantifying uncertainties due to chemistry modelling-evaluation of tropospheric composition
- simulations in the CAMS model (cycle 43r1)," Geoscientific Model Development, vol. 12, no. 4, pp. 1725–1752, 2019.
- 796