

# Quantifying NO<sub>x</sub> emissions in Egypt using TROPOMI observations

Anthony Rey-Pommier<sup>1</sup>, Frédéric Chevallier<sup>1</sup>, Philippe Ciais<sup>1</sup>, Grégoire Broquet<sup>1</sup>, Theodoros Christoudias<sup>2</sup>, Jonilda Kushta<sup>2</sup>, Didier Hauglustaine<sup>1</sup> and Jean Sciare<sup>2</sup>.

<sup>1</sup> Laboratoire des Sciences du Climat et de l'Environnement, LSCE/IPSL, CEA-CNRS-UVSQ, Université Paris-Saclay, 91190 Gif-sur-Yvette, France

<sup>2</sup> The Cyprus Institute, Climate and Atmosphere Research Center, 2121 Nicosia, Cyprus

**Correspondence:** Anthony Rey-Pommier (anthony.rey-pommier@lsce.ipsl.fr)

**Abstract.** Urban areas and industrial facilities, which concentrate the majority of human activity and industrial production, are major sources of air pollutants, with serious implications for human health and global climate. For most of these pollutants, emission inventories are often highly uncertain, especially in developing countries. Spaceborne measurements from the TROPOMI instrument, onboard the Sentinel-5 Precursor satellite, are used to retrieve nitrogen dioxide (NO<sub>2</sub>) column densities at high spatial resolution. Here, we use two years of TROPOMI retrievals to map nitrogen oxides (NO<sub>x</sub> = NO + NO<sub>2</sub>) emissions in Egypt with a top-down approach using the continuity equation in steady state. Emissions are expressed as the sum of a transport term and a sink term representing the three-body reaction comprising NO<sub>2</sub> and hydroxyl radical (OH). This sink term requires information on the lifetime of NO<sub>2</sub>, which is calculated with the use of the CAMS near-real-time temperature and OH concentration fields. We compare this derived lifetime with the lifetime inferred from the fitting of NO<sub>2</sub> line density profiles in large plumes with an exponentially modified Gaussian function. This comparison, which is conducted for different samples of NO<sub>2</sub> patterns above the city of Riyadh, provides information on the reliability of the CAMS near-real-time OH concentration fields; it also provides some hint on the vertical levels that best represent typical pollution sources in industrial areas and megacities in the Middle East region. In Egypt, total emissions of NO<sub>x</sub> are dominated by the sink term, but they can be locally dominated by wind transport, especially along the Nile where human activities are concentrated. Megacities and industrial regions clearly appear as the largest sources of NO<sub>x</sub> emissions in the country. Our top-down model infers emissions with a marked annual variability. By looking at the spatial distribution of emissions at the scale of different cities with different industrial characteristics, it appears that this variability is consistent with national electricity consumption. We detect lower emissions on Fridays, which are inherent to the social norm of the country, and quantify the drop in emissions in 2020 due to the COVID-19 pandemic. Overall, our estimations of NO<sub>x</sub> emissions for Egypt are 7.0% higher than the CAMS-GLOB-ANT\_v4.2 inventory, and significantly differ in terms of seasonality.

## 1 Introduction

Economic growth in developing countries has led to a strong increase of urban air pollution (Baklanov et al., 2016). Among the different pollutants, nitrogen oxides are key species. They are generally the products of fuel combustion, such as the burning of hydrocarbons in the air at high temperature. The main sources of these compounds are vehicle engines, but also heavy industrial facilities such as power plants, iron and steel mills (Tang et al., 2020) and cement kilns (Kim et al., 2020). Their accumulation in the lowest layers of the troposphere contributes to the formation of smog and acid rain (Singh and Agrawal, 2007). They also have a significant effect on human health, as they can cause various respiratory diseases (EPA, 2016). To deal with these phenomena, national and regional governments generally enact a series of air pollution control strategies, which typically take the form of bans on certain polluting technologies, with the aim of reducing the concentration of pollutants at the local level to targets that must be achieved within a given timeframe. These strategies, which also help driving technological innovation, have had a significant effect in Europe (Crippa et al., 2016).

In Egypt, population growth, urbanisation, socio-economic development and the associated increase in the vehicle fleet led to a major degradation of air quality in the last decades, especially in highly populated areas such as Greater Cairo and the Nile Delta (El-Magd et al., 2020) which gather the majority of the population. The Ministry of State for the Environment has thus initiated new policies that aim to reduce pollution levels throughout the country, through technical mitigation of emissions, emission standards for vehicles and intersectoral collaboration (UNEP,

2015). However, Egypt, like most developing countries, lacks the local infrastructure to access detailed information on technical factors such as energy consumption or fuel type and technology, leading to discrepancies in inventories (Xue and Ren, 2012). As a consequence, the monitoring of emissions, which is important to evaluate the effects of air pollution control policies, is of limited reliability.

To overcome the uncertainties in the emission inventories, the use of independent observations systems is becoming increasingly prevalent. In this study, we investigate the use of satellite remote sensing of atmospheric concentrations to improve the quantification of  $\text{NO}_x$  emissions in Egypt. Spectrally resolved satellite measurements of solar backscattered radiation enable the quantification of  $\text{NO}_2$  and other trace gases absorbing in the UV-Visible spectral range based on their characteristic spectral absorption patterns. Tropospheric vertical column densities, i.e. vertically integrated  $\text{NO}_2$  concentrations in the troposphere, have been providing information on the spatial distribution of tropospheric  $\text{NO}_2$  at global scale for nearly 30 years, allowing the identification of different sources of  $\text{NO}_x$  and the quantification of the associated emissions (Leue et al., 2001; Martin et al., 2003; Mijling and Van Der A, 2012; de Foy et al., 2015; Goldberg et al., 2019; Beirle et al., 2019; Lorente et al., 2019; Lange et al., 2021). In October 2017, the Sentinel-5 Precursor satellite was launched. Its main instrument is the TROPOspheric Monitoring Instrument (TROPOMI), which provides tropospheric  $\text{NO}_2$  column densities at high spatial resolution with a large swath width and with a daily frequency (Veefkind et al., 2012). By applying the steady-state continuity equation (Beirle et al., 2019; Lama et al., 2020), it is possible to build a top-down model that directly quantifies  $\text{NO}_x$  emissions from these  $\text{NO}_2$  column densities, provided that some key parameters (wind, temperature, hydroxyl radical concentration and concentration ratio between  $\text{NO}_x$  and  $\text{NO}_2$ ) are correctly estimated. This model is used to quantify the anthropogenic  $\text{NO}_x$  emissions in Egypt for a 2-year period, from November 2018 to November 2020.

This paper is organised as follows: Section 2 provides a description of the datasets used in this study. Section 3 explains the build-up and the limits of the top-down approach used to quantify  $\text{NO}_x$  emissions in Egypt. It also presents a method for validating the model parameters by using  $\text{NO}_2$  line density profiles over Riyadh, Saudi Arabia. Section 4 presents the analysis of this validation method. It presents the location of the main  $\text{NO}_x$  sources in Egypt and evaluates the vertical sensitivity of the model. It also assesses the ability of the model to show less human activity on Fridays and during the lockdown that took place during the COVID-19 pandemic. It finally confronts the inferred emissions with different inventories in terms of amplitude and seasonality. Section 5 presents our conclusion and general remarks.

## 2 Instrumentation and data

### 2.1 TROPOMI $\text{NO}_2$ retrievals

The TROPOspheric Atmosphere Monitoring Instrument (TROPOMI), onboard the European Space Agency’s (ESA) Sentinel-5 Precursor (S-5P) satellite, provides measurements for atmospheric composition. TROPOMI is a spectrometer observing wavelengths in the infrared, visible and ultraviolet light at around 13:30 local time. The UV-Visible spectral band at 405-465 nm is used to retrieve  $\text{NO}_2$ . Other spectral bands are used to observe methane, formaldehyde, sulphur dioxide, carbon monoxide and ozone, as well as aerosols and cloud physical properties. The very high spatial resolution offered by TROPOMI (originally  $3.5 \times 7 \text{ km}^2$  at nadir, improved to  $3.5 \times 5.5 \text{ km}^2$  since 6 August 2019) provides unprecedented information on tropospheric  $\text{NO}_2$ . Its large swath width ( $\sim 2600 \text{ km}$ ) makes it possible to construct  $\text{NO}_2$  images on large spatial scales. Those images greatly improve the potential for detecting highly localised pollution plumes above the ground, identifying small-scale emission sources but also estimating emissions from megacities, industrial facilities and biomass burning. We use TROPOMI  $\text{NO}_2$  retrievals (L2 data, OFFL stream, product version 1.0.0 and 1.1.0 successively) from November 2018 to November 2020 over Egypt. We also use them over Saudi Arabia, and more specifically over the city of Riyadh, to evaluate the reliability of other parameters. This will be explained in Section 3.3. Both countries have an arid climate, which offers a large number of clear-sky days throughout the year, enabling the calculation of monthly averages based on more than 20 observations. They are also the host to many large plumes of pollutants due to high human concentrations along rivers and around megacities, which allows us to observe high  $\text{NO}_2$  concentration patterns with a high signal-to-noise ratio. TROPOMI products provide a quality assurance value  $q_a$ , which ranges from 0 (no data) to 1 (high-quality data). For our analysis of concentrations, we selected  $\text{NO}_2$  retrievals with  $q_a$  values greater than 0.75, which systematically correspond to clear-sky conditions (Eskes et al., 2019). TROPOMI soundings are gridded at a spatial resolution of  $0.1^\circ \times 0.1^\circ$  with daily coverage. This resolution is lower than that of the instrument; the gridding thus provides a grid for which most  $\text{NO}_2$  columns correspond to one or more measurements. The observed plumes remain correctly resolved. Cells without measurements are infrequent, which facilitates the calculation of derivatives.

## 2.2 Wind data

100 The horizontal wind  $\mathbf{w} = (u, v)$  is taken from the European Centre for Medium-Range Weather Forecasts (ECMWF) ERA5 data archive (fifth generation of atmospheric reanalyses) at a horizontal resolution of  $0.25^\circ \times 0.25^\circ$  on 37 pressure levels (Hersbach et al., 2020). The hourly values have been linearly interpolated to the TROPOMI orbit timestamp and re-gridded to a  $0.1^\circ \times 0.1^\circ$  resolution.

## 2.3 CAMS real-time fields

105 The Copernicus Atmospheric Monitoring Service (CAMS) global near-real-time service provides analyses and forecasts for reactive gases, greenhouse gases and aerosols on 25 pressure levels with a horizontal resolution of  $0.4^\circ \times 0.4^\circ$  and a temporal resolution of 3 hours (Huijnen et al., 2016). For the calculation of  $\text{NO}_x$  emissions from TROPOMI observations, we use CAMS concentration fields of nitrogen oxides (NO and  $\text{NO}_2$ ) and hydroxyl radical (OH). We also use the CAMS temperature field  $T$ . NO and  $\text{NO}_2$  concentrations are used to account for chemical processes  
110 that take place in polluted air. Anthropogenic activities produce mainly NO, which is transformed into  $\text{NO}_2$  by reaction with ozone  $\text{O}_3$ .  $\text{NO}_2$  is then photolyzed during the day, reforming NO (Seinfeld, 1989). This photochemical equilibrium between NO and  $\text{NO}_2$  can be highlighted with the  $\text{NO}_x:\text{NO}_2$  concentration ratio, whose value depends on local conditions, allowing to perform a conversion from  $\text{NO}_2$  production to  $\text{NO}_x$  emissions. The reason for the use of OH is different. OH is the main oxidant that controls the ability of the atmosphere to remove pollutants such  
115 as  $\text{NO}_2$  (Logan et al., 1981). It is mainly produced during daylight hours by interaction between water and atomic oxygen produced by ozone dissociation (Levy, 1971). In air that is directly influenced by pollution, another source of OH is due to a reaction between NO and  $\text{HO}_2$ . This reaction, referred to as the  $\text{NO}_x$  recycling mechanism, illustrates the nonlinear dependence of the OH concentration on  $\text{NO}_2$  (Valin et al., 2011; Lelieveld et al., 2016). Since the OH lifetime is typically of less than a second, its concentration in the troposphere is very low and difficult to measure. As  
120 a consequence, global analyses, which estimate OH concentrations from other variable species (Li et al., 2018; Wolfe et al., 2019), provide a representation for OH concentrations with high associated uncertainties. Therefore, the CAMS OH concentrations are used here to account for the  $\text{NO}_2$  oxidation to form nitric acid ( $\text{HNO}_3$ ), whose representation is explained in Section 3.1. Finally, the temperature field is used to control variations in the kinetic parameters (Burkholder et al., 2020). The hourly values are also linearly interpolated to the TROPOMI orbit timestamp and  
125 re-gridded to a  $0.1^\circ \times 0.1^\circ$  resolution.

## 2.4 Background removal

Detecting traces of anthropogenic emissions in TROPOMI  $\text{NO}_2$  images is not a straightforward process. The  $\text{NO}_2$  signal from a sparsely populated area or a small industrial facility may be covered by numerical noise or by the signal generated by natural  $\text{NO}_x$  emissions. In the absence of anthropogenic sources, TROPOMI observes  $\text{NO}_2$   
130 concentrations which constitute a tropospheric background of  $\sim 0.5 \times 10^{15}$  molecules. $\text{cm}^{-2}$ . At the global scale, this background is the result of different sources. In the lower troposphere, natural  $\text{NO}_x$  emissions are mostly due to fires and soil emissions (Yienger and Levy, 1995; Hoelzemann et al., 2004). In the upper troposphere however, sources include lightning, convective injection and downwelling from the stratosphere (Ehhalt et al., 1992), but the factors controlling the resulting concentrations are poorly understood. According to state-of-art estimates, anthropogenic  
135  $\text{NO}_x$  accounts for most of the emissions at the global scale, whereas natural emissions from fires, soils and lightning are less significant at the global scale and do not exceed a share of 35% combined (Jaeglé et al., 2005; Müller and Stavrou, 2005), although associated errors can be very high. In eastern China, the non-anthropogenic share of total  $\text{NO}_x$  emissions is variable but does not exceed 20% (Lin, 2012). Egypt being a desertic region and not very conducive to lightning, we expect the share of those non-anthropogenic emissions to be smaller. To estimate anthropogenic  $\text{NO}_x$   
140 emissions, it is therefore necessary to remove this share.

With an atmospheric lifetime of about a few hours, the presence of  $\text{NO}_2$  is relatively short. Consequently, the majority of  $\text{NO}_2$  is not transported far downwind from its sources. Thus, near-surface  $\text{NO}_2$  concentrations are generally high over industrial facilities and densely populated areas that need to be identified. Because Egypt's population is almost entirely located along the Nile River and its delta, the study of  $\text{NO}_x$  emissions in this country cannot therefore  
145 be reduced to the study of a small number of point sources, as it would be the case for several other parts of the Middle East region, and must be carried out in the form of a mapping of the country. Further discussion is provided in Section 3.3. To identify urban areas in Egypt, we use the Socioeconomic Data and Applications Center (SEDAC) GRUMP (Global Rural-Urban Mapping Project) data archive, which comprises eight global datasets, including a population density grid provided at a resolution of 30 arc seconds, with population estimates normalised for the year  
150 2000 (CIESIN, 2019). We combine this database with field data giving the location of industrial facilities from energy-

intensive industries in the region (data have been retrieved from the Global Energy Observatory for oil and gas-fired power plants, from IndustryAbout for aluminium and iron smelters, from the work of Elvidge et al. (2016) for flaring sites, and from the work of Steven J. Davis and Dan Tong for cement plants; links at the end of this article).

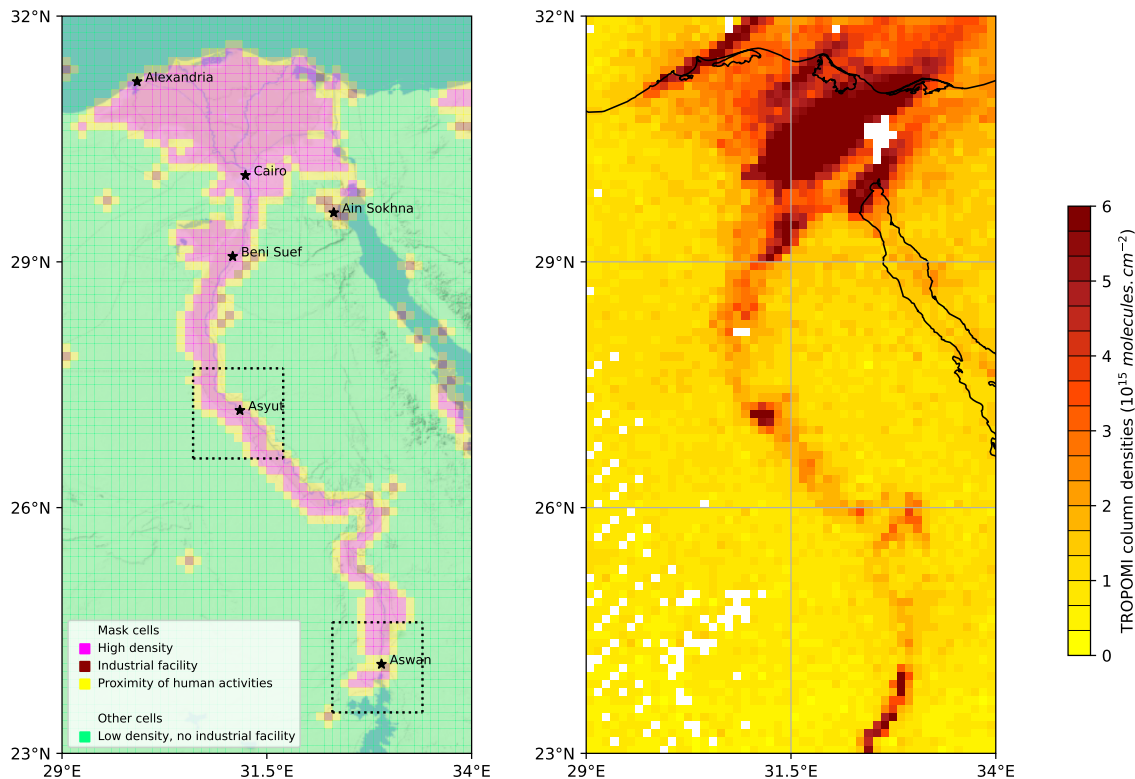


Figure 1: (left) Part of Egypt centered on Nile River. Within this domain, pink cells represent locations with an average human density above  $100 \text{ hab.km}^{-2}$ , brown cells represent locations with industrial facilities outside cities, and yellow cells represent locations in their vicinity. These cells constitute the mask used to calculate anthropogenic emissions. Outside this mask, green cells represent areas which do not correspond to any of the three criteria, considered to be void of human activity. Five large cities in the country and the industrial area of Ain Sokhna are denoted with stars. Two smaller domains centered around the cities of Asyut and Aswan are represented with dotted lines; their use is presented in Section 4.6. (right) TROPOMI observation of  $\text{NO}_2$  slant column densities above Nile valley on 3 January 2019. White pixels correspond to areas with low quality data ( $q_a < 0.75$ ) or no data.

These datasets are used to remove the non-anthropogenic part of the  $\text{NO}_x$  emissions signal. We conduct this removal by subtracting the mean emissions over areas without human activity from the mean emissions over industrial and densely populated areas. In order to perform this distinction between these two types of areas, our study is carried out using a mask within a  $0.1^\circ \times 0.1^\circ$  grid. A grid cell is considered to be part of the mask if it has a population density higher than a threshold of  $100 \text{ hab.km}^{-2}$ , or if its centre is close to an industrial facility. Otherwise, the cell is considered to be part of the "background", i.e. outside the mask. In order to avoid any smearing that would correspond to abnormally high emissions outside urban and industrial centres (which can happen if the wind is poorly estimated), transition cells (in the immediate vicinity of the mentioned mask cells) are also considered to be mask cells. Figure 1 shows the distinction between mask cells and background cells on our domain in Egypt that lies between parallels  $23^\circ\text{N}$  and  $32^\circ\text{N}$  and meridians  $29^\circ\text{E}$  and  $34^\circ\text{E}$ . Most of the mask cells are located in the Nile area. Some mask cells are also found on the coast or in isolated parts in the desert. They correspond to remote industrial facilities, including major flaring sites, or sparsely populated industrial centres such as Ain Sokhna's industrial area. The domain comprises  $n_m = 949$  mask cells and  $n_b = 3692$  background cells. The mathematical description of the background removal is outlined in Section 3.4.

## 2.5 Emission inventories

We compare TROPOMI-derived  $\text{NO}_x$  emissions to the Emissions Database for Global Atmospheric Research (EDGARv5.0) for 2020 and the CAMS global anthropogenic emissions (CAMS-GLOB-ANT\_v4.2) inventory released in 2020. Both datasets provide  $0.1^\circ \times 0.1^\circ$  gridded emissions for different sectors on a monthly basis. EDGARv5.0

emissions are based on activity data (population, energy production, fossil fuel extraction, industrial processes, agricultural statistics, etc.) derived from the International Energy Agency (IEA) and the Food and Agriculture Organization (FAO), corresponding emission factors, national and regional information on technology mix data and end-of-pipe measurements. The inventory covers the years 1970-2015 and differs from the previous version EDGARv4.3.2 which does not use splitting factors derived from the Energy Information Administration (EIA) data on fuel consumption of coal, oil and natural gas for specific countries (Crippa et al., 2020). CAMS-GLOB-ANT\_v4.2 is developed within the framework of the Copernicus Atmospheric Monitoring Service (Granier et al., 2019). For this inventory, NO<sub>x</sub> emissions are based on various existing sectors in the EDGARv4.3.2 emissions from 2000-2012 which are used as a basis for 2010 emissions and are extrapolated to the current year using 2011-2014 sector-based trends from the Community Emissions Data System (CEDS) inventory (Hoesly et al., 2018). From one inventory to another, the names and definitions of the sectors may vary. In EDGARv5.0 and CAMS-GLOB-ANT\_v4.2, the emissions for a given country are derived from the type of technologies used, the dependence of emission factors on fuel type, combustion conditions, as well as activity data and low resolution emission factors (Janssens-Maenhout et al., 2019).

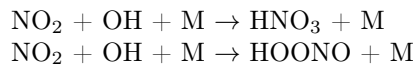
### 3 Method

#### 3.1 Calculation of NO<sub>2</sub> production from TROPOMI observations

As a first step, we use tropospheric NO<sub>2</sub> vertical column densities  $\Omega_{\text{NO}_2}$  to derive top-down NO<sub>2</sub> production maps. Vertical column densities  $\Omega_{\text{NO}_2}$  are obtained from TROPOMI slant column densities using an air mass factor (AMF) which is also provided by TROPOMI. Previous studies have shown that the use of the AMF is a source of structural uncertainty in NO<sub>2</sub> retrievals (Boersma et al., 2004; Lorente et al., 2017). In polluted environments, this source of uncertainty becomes non-negligible. Here, the AMF does not vary much temporally throughout the studied period, and is around 1.6 for mask cells and around 1.9 for background cells. The difference between the two types of cells is probably due to a different albedo between the urban environment and desert areas. Using the horizontal wind  $\mathbf{w}$ , the NO<sub>2</sub> flux is given as  $\Omega_{\text{NO}_2} \mathbf{w}$ . The divergence of this flux can be added to the local time derivative  $\frac{\partial \Omega_{\text{NO}_2}}{\partial t}$  to balance NO<sub>2</sub> sources  $e_{\text{NO}_2}$  and sinks  $s_{\text{NO}_2}$  according to the continuity equation:

$$\frac{\partial \Omega_{\text{NO}_2}}{\partial t} + \text{div}(\Omega_{\text{NO}_2} \mathbf{w}) = e_{\text{NO}_2} - s_{\text{NO}_2} \quad (1)$$

In steady state, the time derivative disappears and the mass balance is reduced to three terms. The NO<sub>2</sub> production can thus be estimated by taking the combined effect of atmospheric transport losses and the different sinks. For the transport term, we calculate numerical derivatives with a fourth-order central-finite difference scheme for each cell of the domain. Moreover, since the local overpass time of TROPOMI occurs in the middle of the day, NO<sub>2</sub> losses are largely dominated by the three-body reaction involving NO<sub>2</sub> and OH (Seinfeld, 1989). Two channels have been identified for this reaction (Burkholder et al., 2020), leading to the production of nitric acid HNO<sub>3</sub> and pernitrous acid HOONO:



For the OH concentrations that are considered in this region ( $1 - 20 \times 10^6$  molecules.cm<sup>-3</sup>), the reactions above follow first order kinetics. The total sink term can therefore be calculated as  $s_{\text{NO}_2} = \Omega_{\text{NO}_2} / \tau$  with:

$$\tau = \frac{1}{k_{\text{mean}}(T, [\text{M}]) \cdot [\text{OH}]} \quad (2)$$

$\tau$  appears here as the characteristic mixed lifetime of NO<sub>2</sub> in the atmosphere. The reaction rate  $k_{\text{mean}}$  characterises the reactions between NO<sub>2</sub> and OH and depends on atmospheric conditions. (Burkholder et al., 2020) provide a general expression of this rate as a function of both temperature  $T$  and total air concentration  $[\text{M}]$ . Note that HOONO can be rapidly decomposed back to NO<sub>2</sub> and OH in the lower troposphere. We assume here that this decomposition is slow and does not affect the NO<sub>2</sub> horizontal gradients. Both pathways are therefore taken into account, and the value of  $k_{\text{mean}}$  represent the total loss of NO<sub>2</sub> due to OH, with a contribution of the HOONO forming reaction between 5 to 15% under atmospheric conditions (Sander et al., 2011; Nault et al., 2016). Thus, the NO<sub>2</sub> production can be calculated as the sum of a transport term and a sink term:

$$e_{\text{NO}_2} = \text{div}(\Omega_{\text{NO}_2} \mathbf{w}) + \Omega_{\text{NO}_2} / \tau \quad (3)$$

The treatment for NO<sub>x</sub> removal is simplified here. NO<sub>x</sub> concentrations are influenced by other sinks. (Stavrakou et al., 2013) showed that the reaction between NO<sub>2</sub> and OH forming HNO<sub>3</sub> accounted for most of total NO<sub>x</sub> loss at

the global scale, but with high uncertainties associated with other sinks. Here, the features of the climate in Egypt during daytime hinder many processes to have a significant effect. The following  $\text{NO}_x$  sinks, which can be of notable importance at the global scale, are not taken into account here:

220 -  $\text{NO}_2$  deposition through the leaf stomata of vegetation. This sink can be significant in forested areas. In Egypt, the leaf area index is very low, except in the croplands of the Nile Delta where it is comparable to that of southern Europe or the western United States (Fang et al., 2019), for which the corresponding lifetime was of about 10-100 h (Delaria et al., 2020), i.e. about an order of magnitude larger than the lifetimes calculated here. To our knowledge, there are no studies focusing on the corresponding lifetimes for croplands, and we therefore do not take this sink into  
225 account.

-  $\text{NO}_x$  oxidation by organic radicals to produce alkyl and multifunctional nitrates (Sobanski et al., 2017). This sink increases with the concentration of volatile organic compounds (VOC), whose presence cannot be excluded in Egypt. Different models have estimated low biogenic isoprene emissions in the region (Wiedinmyer et al., 2006; Guenther et al., 2006). These emissions are concentrated around the Nile River and its delta, and do not exceed  $15 \text{ mg.m}^{-2}.\text{day}^{-1}$ .  
230 They are certainly noticeable and higher in summer than in winter, and contrast with the rest of the country, but they remain low compared to other regions in the world. They are, for instance, about an order of magnitude lower than in the forested areas of the eastern US, where the corresponding sink accounts for between 30% and 60% of the total  $\text{NO}_x$  sink (Romer Present et al., 2020). Furthermore, at large  $\text{NO}_2$  concentrations (compared to VOC concentrations), the share of this sink in the total  $\text{NO}_x$  loss is weakened compared to that of  $\text{HNO}_3$  (Romer Present et al., 2020). The effect  
235 of biogenic emissions of VOC can therefore be considered minor. However, VOC emissions can also be of anthropogenic origin, especially in urban areas, where they are difficult to estimate. To our knowledge, there is no study evaluating the competition of the two sinks in Egypt or in a region with similar features and we therefore do not account for this reaction in our calculations.

- NO reaction with  $\text{HO}_2$  to produce nitric acid (Butkovskaya et al., 2005), whose yield is strongly enhanced in  
240 presence of water vapour (Butkovskaya et al., 2009). Here, we neglect this reaction as the corresponding reaction rate is lower by a factor 3 to 8 in dry conditions (Butkovskaya et al., 2005);

- NO conversion to  $\text{NO}_3$ , the latter being in thermal equilibrium with  $\text{NO}_2$  and  $\text{N}_2\text{O}_5$ . This sink, which takes place via heterogeneous processes, has a significant contribution during nighttime in the Mediterranean region (Friedrich et al., 2021), is neglected at 13:30 when OH is close to its daily maximum;

245 -  $\text{NO}_2$  reversible reaction with peroxyacetyl radical to produce peroxyacetylnitrate (Moxim et al., 1996). In the Nile Delta region, PAN concentrations in the lower troposphere are significantly below the global average (Fischer et al., 2014), possibly due to high temperatures favoring short PAN lifetimes. Moreover, its production peaks in the late afternoon and early evening (Seinfeld, 1989). We therefore do not consider this sink in the representation of  $\text{NO}_x$  emissions at 13:30;

250 -  $\text{NO}_2$  uptake onto black carbon particles (Longfellow et al., 1999). This uptake is of limited amount in the Mediterranean region (Friedrich et al., 2021).

All these processes not being accounted for, the reaction between  $\text{NO}_2$  and OH is the only sink that is considered in our calculations to provide an indication of  $\text{NO}_x$  emissions. Section 4.7 details the consequences of not considering these various minor sinks on the results.

### 255 3.2 Interpolation to daily average emissions

All parameters are evaluated at 13:30 local time, which means that the  $\text{NO}_2$  production is calculated at the same moment. In Egypt, the maximum and minimum electricity consumption are reached around 20:00 and 6:00 local time respectively, and inter-daily consumption differences have been weakened by the increasing sales of air conditioning and ventilation systems in the past decades (Attia et al., 2012). The daily load profiles provided by the National  
260 Egyptian Electricity Holding Company show that the mean daily electricity consumption corresponds approximately to the consumption at 13:30 in the country (EEHC, 2021). The difference between the two quantities being small both in summer (about +2 to -3%) and winter (about -2 to -6%), we consider our inferred emissions as representative of the average activity in Egypt at any time of the year. This assumes that electricity consumption dominates the emissions of the country, or that the other emitting sectors have a similar daily profile. This can be justified.  
265 According to CAMS-GLOB-ANT\_v4.2, the power sector accounts for 50 to 60% of total  $\text{NO}_x$  emissions in Egypt. EDGARv5.0 presents a lower share (40 to 45% of total emissions). Moreover, for both inventories, the transport sector accounts for the majority of the remaining emissions. According to the traffic congestion index in Cairo ([https://www.tomtom.com/en\\_gb/traffic-index/cairo-traffic/](https://www.tomtom.com/en_gb/traffic-index/cairo-traffic/)), the congestion level around 13:30 seems to be slightly higher than during the morning peak, but lower than the during night peak. Traffic emissions at this moment of the  
270 day could therefore be representative of the average traffic emissions as well.

### 3.3 Validation of CAMS OH concentration using line density calculations for Riyadh

When the transport term is integrated over large spatial scales, it cancels out due to the mass balance in the continuity equation between NO<sub>2</sub> sources and NO<sub>2</sub> sinks. At first order, the integration of the inferred emissions over the whole domain (of about 490,000 km<sup>2</sup>) thus reflects chemical losses of the sink term. In this term, the NO<sub>2</sub> lifetime calculation involves the reaction rate  $k_{\text{mean}}$ , whose annual variability is low due to small changes in Egyptian midday temperatures throughout the year, and OH concentration, whose annual variability is highly marked. In Egypt, tropospheric OH concentrations, which are strongly correlated with solar ultraviolet radiation (Rohrer and Berresheim, 2006) and NO<sub>x</sub> emissions, are higher in summer than in winter. To ensure an appropriate representation of the OH field by CAMS data, we select a large number of TROPOMI images characterised by a homogeneous wind field, in which we calculate the NO<sub>2</sub> lifetime according to Equation (2), where [OH] corresponds to the near-real-time CAMS data and  $k_{\text{mean}}$  is calculated with the formula from Burkholder et al. (2020). We compare this value with the lifetime determined by a method initially developed by Beirle et al. (2011), and expanded by Valin et al. (2013) by introducing a rotation of the image to refine the chemical lifetime. This method consists in fitting an exponentially modified Gaussian function (EMG) to NO<sub>2</sub> line density profiles. These profiles correspond to the integrated NO<sub>2</sub> columns along the mean wind direction in the pollution pattern and centered around the source. They are obtained by rotating TROPOMI images in the mean wind direction and using the values of the nearest columns in a 100 km<sup>2</sup> area. Line density profiles are generated on a span of 300 km. An example is given in Figure 3. Within the average profile, the NO<sub>2</sub> burden and lifetime can be derived from the parameters that describe the best statistical fit. The EMG model is expressed as follows (Lange et al., 2021):

$$\langle \Omega_{\text{NO}_2} \rangle(x|B, A, x_0, \mu, \sigma) = B + \frac{A}{2x_0} \exp\left(\frac{\mu - x}{x_0} + \frac{\sigma^2}{2x_0^2}\right) \operatorname{erfc}\left(-\frac{1}{\sqrt{2}}\left(\frac{x - \mu}{\sigma} - \frac{\sigma}{x_0}\right)\right) \quad (4)$$

Here,  $x$  is the distance in the downwind-upwind direction,  $B$  is the NO<sub>2</sub> background,  $A$  is the total number of NO<sub>2</sub> molecules observed in the vicinity of the point source,  $x_0$  is the e-folding distance downwind, representing the exponential length scale of NO<sub>2</sub> decay,  $\mu$  is the location of the apparent source relative to the centre of the point source, and  $\sigma$  is the standard deviation of the Gaussian function, representing the length scale of Gaussian smoothing. Using a non-linear least squares fit, we estimate the five unknown parameters:  $A$ ,  $B$ ,  $x_0$ ,  $\mu$  and  $\sigma$ . From the mean wind module  $w_{\text{mean}}$  in the domain, the mean effective NO<sub>2</sub> lifetime  $\tau_{\text{fit}}$  can be estimated using the fitted parameters:

$$\tau_{\text{fit}} = \frac{x_0}{w_{\text{mean}}} \quad (5)$$

The geography of Egypt does not suit the method described here. The Egyptian population is contiguously concentrated along the Nile, which makes it difficult to define point sources isolated from human activity. Furthermore, large isolated cities such as Alexandria or Suez are too close to the coast for the wind to be considered homogeneous. We therefore use the city of Riyadh, Saudi Arabia (24.684°N, 46.720°E) to perform the comparison between the CAMS-induced lifetime and the fit-obtained lifetime. Riyadh has been the focus of anterior studies (Valin et al., 2013; Beirle et al., 2019), and is particularly suitable for several reasons. Firstly, Riyadh is a city within the latitudinal extend of Egypt (1600 km from Cairo) with a climate which is similar to the typical Egyptian climate. Secondly, NO<sub>2</sub> tropospheric columns over Riyadh are high ( $\sim 9 \times 10^{15}$  molecules.cm<sup>-2</sup>), leading to retrievals with a high signal-to-noise ratio. Thirdly, Riyadh is far from the coast, and its flat terrain makes the surrounding wind fields rather homogeneous during most of the year.

As the fitting algorithm is very sensitive to any disturbance that might be induced by NO<sub>2</sub> production from other point sources, it is necessary to identify heavy industrial facilities in the area. Riyadh is also an industrial area, with several power plants located close to the city centre. Figure 2 shows the location of the most important emitters in the region, which include five gas-fired power plants (PP7, PP8, PP9, PP10 and PP14), one oil-fired power plant (PP4) and one cement plant (CP). The five gas power plants, with a total capacity of more than 10 GW, are located in the periphery of the city. These power plants are sufficiently far away from the city centre for TROPOMI to distinguish their own emissions from those of Riyadh's centre with a resolution of  $0.1^\circ \times 0.1^\circ$ , which is not the case for CP and PP4 which are located in the city centre. It is therefore appropriate to restrict the study of NO<sub>2</sub> patterns over Riyadh to days for which the emissions from the city centre and those from the gas power plants do not overlap. This is the case when the wind blows steadily and homogeneously in a north-south direction. Within about 150 km around the city centre, we thus calculate the average wind given by ERA5 and consider the observation as reliable if the mean angle  $\langle \theta \rangle$  of the observations deviates by less than  $40^\circ$  from the north or the south, with a standard deviation  $\sigma_\theta$  of less than  $36^\circ$ . This condition generally leads to a selection of observations with large wind speeds, low winds speeds being often associated with more variable directions. This ensures the NO<sub>2</sub> decay to be dominated by chemical removal

320 and not by the variability of the wind (Valin et al., 2013). Finally, we select observations with clear-sky conditions downstream of the flow (with at least 80% downstream cells with  $q_a > 0.75$ ).

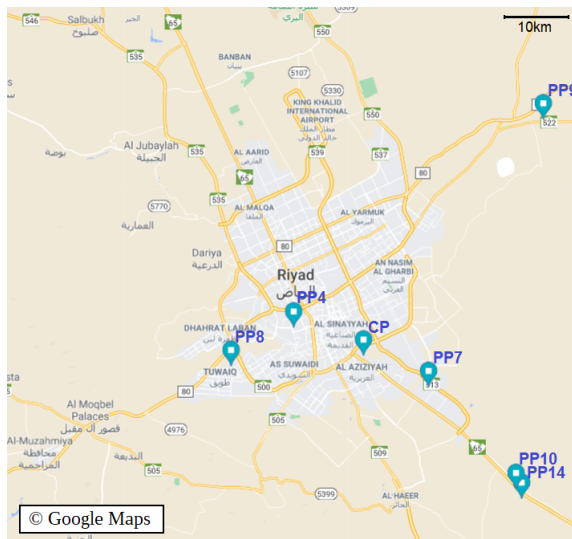


Figure 2: Map of Riyadh's city centre with the surrounding power plants (PP4, 7, 8, 9, 10 and 14) and cement plant (CP). The map has been extracted from © Google Maps.

Our  $0.1^\circ \times 0.1^\circ$  gridding ensures that retrieved lifetimes are governed by physical decay of  $\text{NO}_2$  and not an artifact of the spatial resolution (Valin et al., 2011). The fitting procedure is very sensitive to the wind direction. Instead of manually correcting the ERA5 wind field for individual  $\text{NO}_2$  patterns, the curve fitting is performed for every sample with three different rotation angles, corresponding to the wind direction with a correction of  $-10^\circ$ ,  $0^\circ$  or  $10^\circ$ . A record is kept if one of these three fits leads to a correlation with the corresponding  $\text{NO}_2$  line density whose coefficient is greater than 0.97. Among the remaining samples, we keep those with a value of  $\tau_{\text{fit}}$  greater than 1.0 hr (considered sufficiently high to be relevant). An example of curve fitting is given in Figure 3.

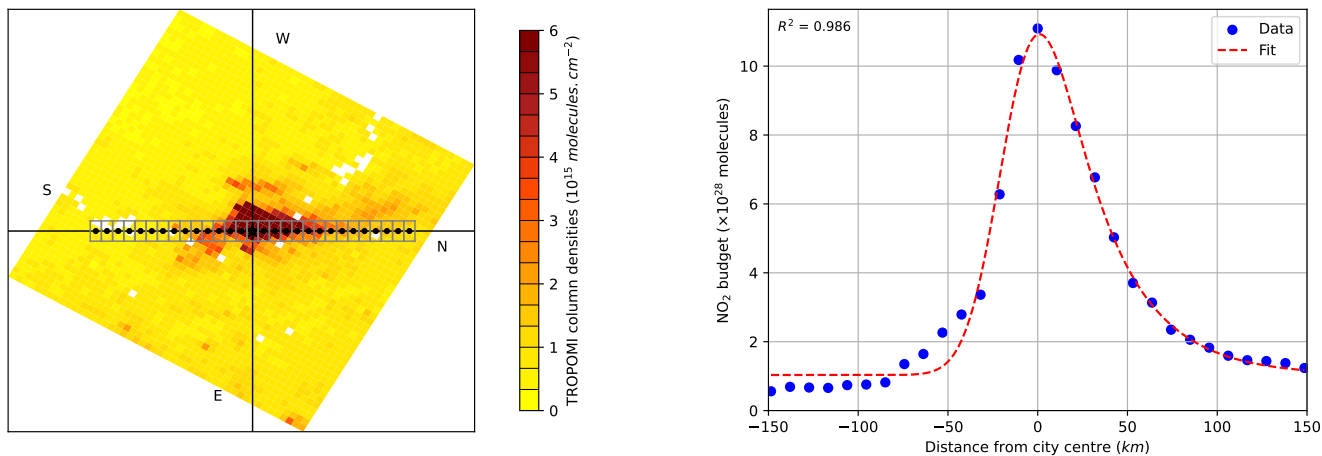


Figure 3: Estimation of the  $\text{NO}_2$  lifetime from a pattern above Riyadh on 11 March 2020: (left)  $\text{NO}_2$  plume rotated with its wind direction around the source (star) to an upwind-downwind pattern. Grey boxes centered around black points indicate the extent of the spatial integration of  $\text{NO}_2$  columns to obtain the  $\text{NO}_2$  line density. Values of cardinal points are noted in black. (right) Corresponding line densities (points) representing the downwind evolution of  $\text{NO}_2$  as function of the distance to Riyadh's city centre, and the corresponding fit according to the exponentially modified Gaussian function (dashed line).

The phenomena under study here take place in the planetary boundary layer (PBL), which in this region has a midday height of about 2 km (Filioglou et al., 2020). TROPOMI observations only provide information on the total  $\text{NO}_2$  content of the tropospheric column, without providing information on the vertical distribution of concentrations. Extracting emissions from concentrations therefore requires a selection on the height at which wind, temperature and OH data are taken. Lama et al. (2020) and Lorente et al. (2019) conducted similar studies using the boundary



layer average wind, while Beirle et al. (2019) chose a vertical level of about 450 m above ground. Because vertical  
 335 transport of  $\text{NO}_x$ , which is emitted mainly from combustion engines and industrial stacks, is generally minor compared  
 to horizontal transport,  $\text{NO}_x$  is confined to the first few hundred metres above ground level. Using PBL-averaged data  
 poses a problem of consistency as wind, temperature and OH concentration values significantly vary within the PBL.  
 As a consequence, we compare the CAMS-induced lifetime  $\tau$  and the fit-induced lifetime  $\tau_{\text{fit}}$  using the parameters ( $\mathbf{w}$ ,  
 [OH] and  $T$ ) at two different vertical levels: a medium level  $\mathcal{A}$  at 925 hPa (about 770 m above ground level), and  
 340 a bottom level  $\mathcal{B}$  at 987.5 hPa (about 210 m). These levels are interpolated from four and two ECMWF or CAMS  
 consecutive pressure levels respectively (1000-850 hPa for level  $\mathcal{A}$  and 1000-975 hPa for level level  $\mathcal{B}$ ). Most mask cells  
 having an altitude between 0 and 150 m, the corresponding pressure variations are small (up to  $\sim 16$  hPa), which  
 allows us to neglect the effects of topography on the position of pressure levels. Figure 4 sums up the selection method  
 for the comparison of lifetimes.

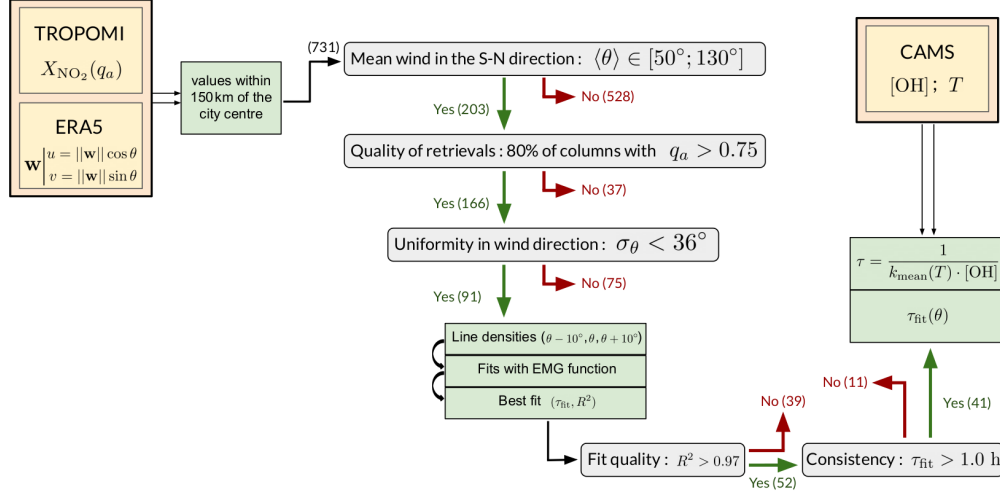


Figure 4: Selection method for  $\text{NO}_2$  patterns over Riyadh. Datasets (yellow-orange) are used to calculate the quantities (light green) that are submitted to different tests (grey). 731 patterns are progressively conserved (green arrows) or rejected (red arrows). At each stage, the number of conserved or rejected patterns are noted within brackets (the value is only given for calculations performed at level  $\mathcal{B}$ ). This selection process compares the lifetimes estimated by the EMG function fitting with TROPOMI line density profiles to the lifetimes calculated according to Equation (2) with CAMS data.

### 3.4 Calculation of anthropogenic $\text{NO}_x$ emissions and comparison with inventories

We calculate  $\text{NO}_x$  emissions on the entire domain from  $\text{NO}_2$  production by using CAMS NO and  $\text{NO}_2$  concentrations. These are not intended to replace TROPOMI observations; they are used to apply the concentration ratio  
 350  $[\text{NO}_x]/[\text{NO}_2] = ([\text{NO}] + [\text{NO}_2])/[\text{NO}_2]$  to account for the conversion of  $\text{NO}_2$  to NO and vice versa. As diurnal NO concentrations in urban areas generally range from 10 to 150 ppb (Khoder, 2009), the characteristic stabilization time  
 of this ratio never exceeds a few minutes (Graedel et al., 1976; Seinfeld and Pandis, 2006). This time being lower than  
 the order of magnitude of the inter-mesh transport time (about 30 min considering the resolution used and the mean  
 wind module in the region), we can reasonably neglect the effect of the stabilization time of the conversion factor on  
 the total composition of the emissions and treat each cell of the grid independently from its neighbours. Beirle et al.  
 (2019) found an annual average of 1.32 for this conversion factor, but CAMS data shows small deviations from this  
 355 value over Egyptian urban areas. We therefore calculate  $\text{NO}_x$  emissions for each cell of the domain as follows:

$$e_{\text{NO}_x} = \frac{[\text{NO}_x]}{[\text{NO}_2]} e_{\text{NO}_2} \quad (6)$$

For convenience, quantities  $\frac{[\text{NO}_x]}{[\text{NO}_2]} \text{div}(\Omega_{\text{NO}_2} \mathbf{w})$  and  $\frac{[\text{NO}_x]}{[\text{NO}_2]} \Omega_{\text{NO}_2} / \tau$  represent the respective contributions of the transport and the sink terms to total  $\text{NO}_x$  emissions. We finally obtain the emissions related to human activity  $E_{\text{NO}_x}$  by removing the arithmetic mean value of  $\text{NO}_x$  emissions above background cells from total emissions:

$$E_{\text{NO}_x} = e_{\text{NO}_x} - \frac{1}{n_b} \sum_{i=1}^{n_b} e_{\text{NO}_x, i} \quad (7)$$

These removed emissions are linked to the  $\text{NO}_2$  background estimated by TROPOMI. This background, which is mostly located in the upper troposphere, is inconsistent with the use of other parameters which are calculated in the lower troposphere. As such, these emissions do not correspond to anthropogenic emissions, but they provide the value of what must be subtracted from the estimates to obtain emissions related to human activity. Such a removal assumes that other processes involved in  $\text{NO}_x$  budgets lead to similar emissions inside and outside the mask, which is not evident, as the majority of background cells are located in the desert or the ocean while the majority of mask cells are located near the Nile River. However, as the processes involved in natural  $\text{NO}_x$  sources lead to emissions much smaller than anthropogenic emissions in polluted areas, we neglect this difference in the following calculations. An alternative would be to calculate an average concentration for the background cells and subtract the corresponding value from the column densities before calculating the emissions. This would pose further reliability problems: for instance, very high  $\text{NO}_2$  concentrations could appear outside the mask due to wind transport (an example is shown on Figure 1). They would lead to an overestimation of the  $\text{NO}_2$  background and thus to an underestimation of the anthropogenic emissions.

Neglecting the part of the country that lies outside the domain, total emissions from the anthropogenic activity of Egypt can then be obtained by integrating  $E_{\text{NO}_x}$  on the whole domain. For robust statistics, these derived emissions can be averaged monthly, enabling a month-by-month comparison with bottom-up inventories. The linearity of the averaging processes ensures the interchangeability of temporal and spatial averages. A monthly average is relevant because it aggregates enough data to limit the impact of the outliers due to uncertainties in wind and OH representation. In addition, it enables the study of monthly  $\text{NO}_x$  emission profiles which reflect changes in human activities throughout the year due to temperature changes, economic constraints and cultural norms.

## 4 Results and discussion

### 4.1 Line densities and $\text{NO}_2$ lifetime

We compare the results of the TROPOMI line densities fits for Riyadh to the lifetime calculated by Equation (2) using CAMS OH data. The two years of TROPOMI observations (from November 2018 to November 2020) provide a wide variety of profiles. For level  $\mathcal{B}$ , Figure 4 also provides the number of samples that are being kept at each stage of the process. Of the 731 observations available, 203 have a wind direction in the cone with a north-south orientation with an aperture of  $40^\circ$  (i.e. between  $340^\circ$  and  $20^\circ$  or between  $160^\circ$  and  $200^\circ$ ). Of the remaining observations, 166 occurred with a sufficiently clear sky to be retained. The criterion of weak variability for the wind direction brings to 91 the number of observations that are kept by the method. On these 91 observations, the line density profiles are calculated and the fits applied. According to Equation (5), the lifetime is calculated using the mean wind module around the point source. The two lifetimes are calculated with the parameters taken at the medium level  $\mathcal{A}$  or at the top level  $\mathcal{B}$ . Of the 91 fits obtained, 51 are of high quality (correlation coefficient between fit function and line density profile greater than 0.97) for level  $\mathcal{A}$  and 52 for level  $\mathcal{B}$ . 39 of these fits lead to a lifetime  $\tau_{\text{fit}}$  greater than 1.0 h for level  $\mathcal{A}$  and 41 for level level  $\mathcal{B}$ . All remaining samples correspond to atmospheric conditions with moderate to fast winds, with a module ranging between 2 and 11 m/s (with an average of 5.9 m/s) for level  $\mathcal{A}$  and between 3 and 8 m/s (with an average of 5.4 m/s) for level  $\mathcal{B}$ . These lifetimes are compared to the corresponding lifetimes obtained from CAMS data in Figure 5, which is divided into seasons for a more convenient comparison. The use of level  $\mathcal{A}$  leads to notable underestimations of the  $\text{NO}_2$  lifetime in autumn compared to the lifetime calculated with the fitting method. This same level leads to lifetime overestimations in winter. This trend is not found with the use of level  $\mathcal{B}$ , which leads to a better reproduction of the lifetimes calculated with the fitting method for the available seasons. Figure 5 shows a linear regression between the two calculated lifetimes. The results are scattered, with a correlation coefficient higher for level  $\mathcal{B}$  (0.408) than for level  $\mathcal{A}$  (0.220). When the intercept of the regression line is forced to zero, the resulting slope is closer to 1 for level  $\mathcal{B}$  (0.998) than for level  $\mathcal{A}$  (1.071).

Although both correlations are weak, level  $\mathcal{B}$  leads to a better match between the lifetime calculated with Equation (2) and the lifetime calculated from line densities. The results that are presented in the following sections (except for Section 4.3) are therefore results of calculations performed with parameters ( $\mathbf{w}$ ,  $[\text{OH}]$ ,  $T$  and  $[\text{NO}_x]/[\text{NO}_2]$ ) estimated at level  $\mathcal{B}$ . Nevertheless, it should be noted that no summer observations were included in the comparison. The main reason for this is the wind direction: of the 188 summer days observed, 178 have a mean wind direction outside the north-south cone over central Riyadh. On the remaining ten days (one for summer 2019 and nine for summer 2020), the ERA5 wind direction is too variable for the fit to be considered relevant, or the fit results in a correlation coefficient below 0.97. It is not clear how correctly the  $\text{NO}_2$  lifetime would be calculated during both summer periods by Equation (2). With OH concentrations being the main driver of this lifetime, we cannot assess the relevance of the

representation of OH concentrations by CAMS data during summer days in the study.

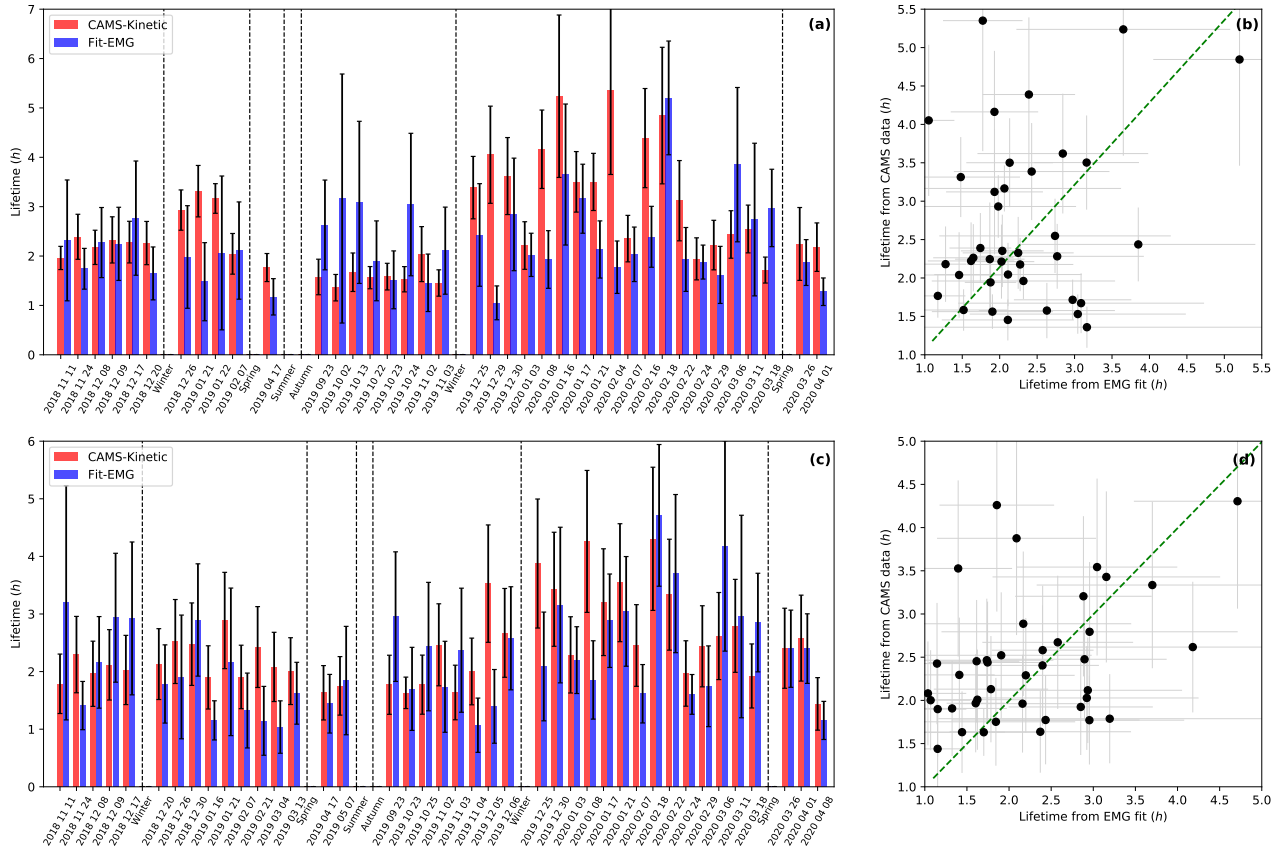


Figure 5: (left) Comparison between CAMS-derived  $\text{NO}_2$  lifetimes and lifetimes from  $\text{NO}_2$  line density fittings with EMG function above Riyadh city centre, for level  $\mathcal{A}$  (a) and  $\mathcal{B}$  (c). The samples presented correspond to patterns in clear-sky conditions for which the mean wind is in the north-south direction with a low variance, and for which the correlation between line density profile and fit gives a correlation coefficient of more than 0.97 and a lifetime of more than 1.0 h.  $\text{NO}_2$  patterns do not have these conditions during the summers of 2019 and 2020. Dashed lines separate the groups of observations by season. (right) Comparison between the two calculated lifetimes for level  $\mathcal{A}$  (b) and  $\mathcal{B}$  (d). A linear regression with an intercept forced to be zero is displayed with a green dashed line.

## 4.2 Mapping of Egypt's $\text{NO}_x$ emissions

As a first step, we try to map  $\text{NO}_x$  emissions in Riyadh using parameters estimated at level  $\mathcal{B}$ . For the period from December 2017 to October 2018 and using a constant lifetime of 4 h, Beirle et al. (2019) estimated at  $6.66 \text{ kg}\cdot\text{s}^{-1}$  the emissions of the corresponding urban area, and a mean rate density of about  $3.7 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  for power plants PP9 and PP10/14, the transport term accounting for about 80 to 90% of this budget. Using the same domain for December 2018 to October 2019 with our method, we found a mean lifetime of 2.94 h and mean emissions of  $5.92 \text{ kg}\cdot\text{s}^{-1}$  for the urban area. We also found smaller rate densities for the power plants (about  $3.4 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  for PP9 and  $3.0 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  for PP10/14), with a smaller contribution of the transport term (about 70%). Despite differences in resolution, AMF calculation, lifetime variability and background removal, the two methods give similar results.

The top-down emission model is then applied to the Egyptian domain with CAMS OH concentration and temperature fields for lifetime calculations. For each cell,  $\text{NO}_x$  emissions are calculated according to Equation (6), resulting in a mapping of Egypt's emissions. The obtained values are averaged monthly from November 2018 to November 2020. Figure 6 shows a composition of the emissions map with the transport term and the sink term for the months of January and July 2019. The corresponding anthropogenic emissions, calculated according to Equation (7), are added. The Nile appears on transport term maps: the divergence calculation complies with what is expected from a line of emitters, i.e. a clear separation of zones of positive divergence from zones of negative divergence with a separation line corresponding to the course of the river. The fact that areas of negative and positive divergence are respectively located to the east and the west of the river is the result of the zonal component of the wind being positive most of the time. Some point sources like Cairo, Alexandria, Asyut or Aswan are easily identifiable. The sink term, directly proportional to the TROPOMI column densities, also highlights these cities. However, unlike the transport term,

which has a similar spatial pattern from month to month, the sink term is clearly stronger in summer than in winter. This is mainly due to a higher lifetime in winter than in summer (4.94 h on average in January 2019 and 2.62 h in July 2019) while the average TROPOMI NO<sub>2</sub> concentrations are slightly higher during winter ( $1.071 \times 10^{15}$  molecules.cm<sup>-2</sup> for January 2019 and  $0.899 \times 10^{15}$  molecules.cm<sup>-2</sup> for July 2019). Over the whole domain, the mean transport term varies throughout the studied period between  $-0.014 \times 10^{15}$  molecules.cm<sup>-2</sup>.h<sup>-1</sup> (December 2019) and  $0.015 \times 10^{15}$  molecules.cm<sup>-2</sup>.h<sup>-1</sup> (May 2020). Thus, it hardly contributes to the NO<sub>x</sub> emission budget, the mean chemical sink term alone varying between  $0.223 \times 10^{15}$  molecules.cm<sup>-2</sup>.h<sup>-1</sup> (December 2019) and  $0.534 \times 10^{15}$  molecules.cm<sup>-2</sup>.h<sup>-1</sup> (September 2020).

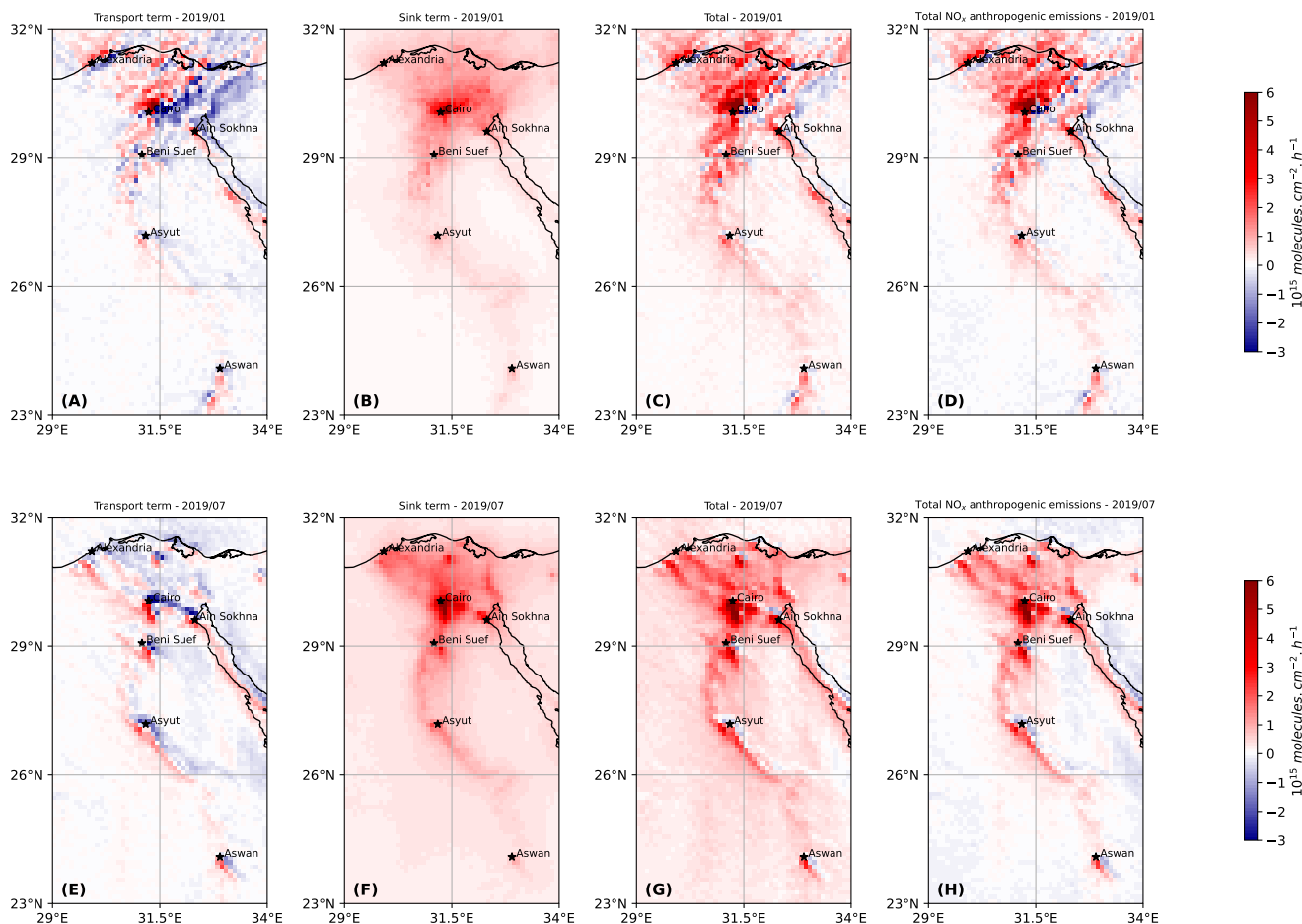


Figure 6: NO<sub>x</sub> emissions above most of Egypt’s territory: (top) transport term (A), sink term (B), resulting emissions (C), and the corresponding anthropogenic emissions after non-anthropogenic background removal (D) for January 2019. (bottom) transport term (E), sink term (F), resulting emissions (G), and the corresponding anthropogenic emissions after background removal (H) for July 2019.

Several cities in the country appear as the main emitters of the country, such as Cairo, Alexandria, Beni Suef, Asyut or Aswan. The industrial area of Ain-Sokhna, located southwest of Suez, also appears as a main emitter. Table 1 compares the monthly values for the sink term and the absolute value of the transport term above five major cities of the country, with populations ranging from 200,000 to 20 million inhabitants, as well as Ain-Sokhna’s area. The mean values for TROPOMI column densities are also provided. According to the results, the capital city of Cairo is by far the largest emitter in the country, largely due to its large population, resulting in high traffic emissions, but also to its intensive industrial activity. Alexandria, the country’s second largest city, is not necessarily the second largest emitter, as its emissions are comparable to those of smaller cities such as Beni Suef or Asyut. However, the three cities concentrate a large amount of industrial activity: Alexandria hosts several oil and gas power plants and a small number of cement factories, while Beni Suef is close to several oil and gas power plants and hosts several flaring sites. Similarly, the city centre of Asyut is close to three oil and gas-fired power plants and a cement factory. This seems to indicate that industrial activity might be the main cause of NO<sub>x</sub> emissions differences between these cities, before population size. This explains why NO<sub>x</sub> emissions from these three cities are comparable to those of the industrial area of Ain Sokhna, which includes several cement facilities, iron smelters and oil and gas plants. It might also explain

why Aswan, which has a population that is comparable to Beni Suef or Asyut, but which does not have any major industrial site, has slightly lower emissions than the two other cities. An additional analysis of the differences between Asyut and Aswan is provided in Section 4.6. Finally, the Gulf of Suez displays relatively large emissions, which might be attributed to the shipping sector, the region being a major gateway for international trade. Because it also hosts several flaring sites, these emissions might also be due to the oil and gas extraction activity.

City	Population density (habitants per square kilometer)	Jan. 2019			Jul. 2019		
		$\Omega_{\text{NO}_2}$	Transport	Sink	$\Omega_{\text{NO}_2}$	Transport	Sink
		( $\mathcal{M}_{\text{NO}_2} \cdot \text{cm}^{-2}$ )	( $\mathcal{M}_{\text{NO}_x} \cdot \text{cm}^{-2} \cdot \text{h}^{-1}$ )		( $\mathcal{M}_{\text{NO}_2} \cdot \text{cm}^{-2}$ )	( $\mathcal{M}_{\text{NO}_x} \cdot \text{cm}^{-2} \cdot \text{h}^{-1}$ )	
Cairo	18,064	9.415	2.903	3.684	5.618	2.022	4.879
Alexandria	9,133	3.034	1.179	0.975	1.674	0.410	1.421
Asyut	1,644	1.708	0.679	0.718	2.137	1.236	1.520
Aswan	319	0.976	0.182	0.473	0.871	0.308	0.523
Beni Suef	2,056	2.950	0.548	1.080	2.321	0.428	1.591
Ain Sokhna	5	3.133	1.256	1.115	2.561	1.346	1.757

Table 1: Comparison between the transport term and the sink term above different cities in Egypt, as well as the industrial region of Ain Sokhna located 45 km southwest of Suez for January and July 2019. TROPOMI vertical  $\text{NO}_2$  columns,  $\text{NO}_x$  emissions and population densities correspond to average values within 18 km from the city centre. Unit  $\mathcal{M}$  stands for a quantity of  $10^{15}$  molecules ( $\text{NO}_2$  or  $\text{NO}_x$ ).

Although these cities and areas can be described as high-emission sites, the term responsible for these emissions differ from one site to the other. Figure 7 shows the contribution of the transport term (taken in absolute value) to total emissions for January and July 2019. Because wind fields are relatively homogeneous along the Nile on spatial scales of less than 100 km,  $\text{NO}_2$  concentration gradients perceived by TROPOMI in the region mainly contribute to the increase of the transport term which can reach similar values as the sink term. However, it is never significantly higher than the sink term: due to a spread of the emissions over large urban areas, the behaviour of these cities is therefore different from that of a point source for which the transport term would be very high (Beirle et al., 2021).

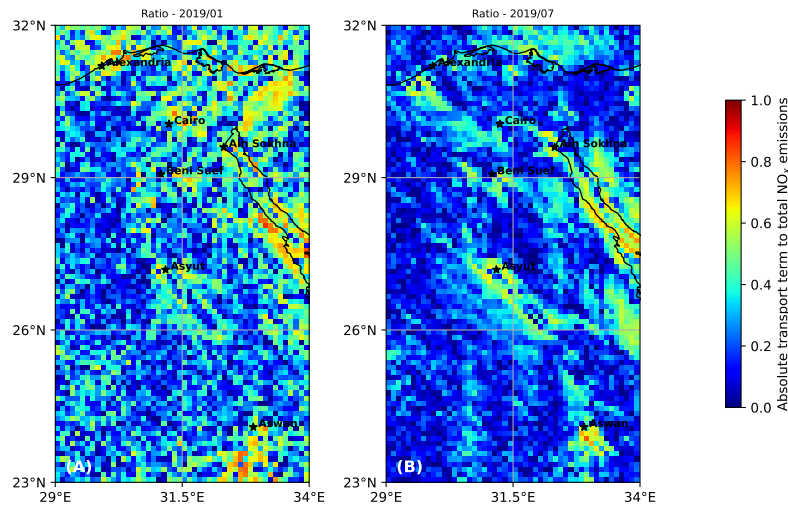


Figure 7: Share of the absolute value of the transport term in the sum of the sink term and the absolute value of the transport term above most of Egypt's territory, indicating the local importance of the transport term in  $\text{NO}_x$  emissions above mask cells. The average of this ratio is shown for January 2019 (A) and July 2019 (B).

Desert areas such as the Libyan Desert, the Eastern Desert and the Sinai region, (located respectively to the west, east and northeast of the Nile) show a very low value for the transport term compared to the sink term, due to the homogeneity of both the wind field and the detected  $\text{NO}_2$  concentrations in these areas. In the case of the Gulf of Suez, the transport term can be 1 to 2 times higher than the sink term, which varies between  $0.4$  and  $1.2 \times 10^{15}$  molecules. $\text{cm}^{-2} \cdot \text{h}^{-1}$ . Those values are slightly higher than the average emissions above background cells areas due to the sink term (about  $0.2 - 0.6 \times 10^{15}$  molecules. $\text{cm}^{-2} \cdot \text{h}^{-1}$ ), but remain quite low compared to the emissions in large cities. This relative predominance of the transport term is explained by a visible gradient of the TROPOMI  $\text{NO}_2$  column densities. The region thus acts as a very thin line of emitters. Nevertheless, this predominance might also be partly due to a poor representation of the wind field. The low resolution of ERA-5 (about 26 km in this region,

475 which is the same order of magnitude as the width of the channel) may misrepresent the wind near the coast, creating artificial gradients.

### 4.3 Vertical analysis

Here we investigate the influence of the choice of the vertical level in the representation of the different model parameters. This influence is of considerable importance, as  $\text{NO}_x$  sources in urban areas can be located at different altitudes. 480 For instance, emissions from the road sector from tailpipes are located at ground level, whereas  $\text{NO}_x$  from power plants and industrial facilities can be emitted from stacks, which are usually located between 50 and 300 m above ground level. Section 4.1 results showed that level  $\mathcal{B}$  was more appropriate for the representation of the  $\text{NO}_2$  lifetime. This level is therefore chosen as a reference for the comparison. We study the effect of a transition from level  $\mathcal{B}$  to level  $\mathcal{A}$  for each of the 3 parameters involved in the representation of the sink term, i.e. temperature  $T$ , hydroxyl radical 485 concentration  $[\text{OH}]$  and concentration ratio  $[\text{NO}_x]/[\text{NO}_2]$ . The results for the averages over mask cells and background cells are given for the months of January, April, July and October 2019 in Table 2. As the wind field is only involved in the transport term whose spatial integration nearly leads to zero, the influence of this parameter is not studied.

		Sink term ( $10^{15} \text{ molecules.cm}^{-2} \cdot \text{h}^{-1}$ )							
level $\mathcal{B}$ (987.5 hPa)	level $\mathcal{A}$ (925 hPa)	Jan. 19 (MASK)	Jan. 19 (BKGD)	Apr. 19 (MASK)	Apr. 19 (BKGD)	Jul. 19 (MASK)	Jul. 19 (BKGD)	Oct. 19 (MASK)	Oct. 19 (BKGD)
$T, [\text{OH}], \frac{[\text{NO}_x]}{[\text{NO}_2]}$	-	0.859	0.253	1.072	0.345	1.125	0.376	0.932	0.277
$[\text{OH}], \frac{[\text{NO}_x]}{[\text{NO}_2]}$	$T$	0.899 (+4.7%)	0.264 (+4.2%)	1.127 (+5.2%)	0.361 (+4.6%)	1.185 (+5.3%)	0.394 (+4.9%)	0.887 (+4.8%)	0.264 (+4.5%)
$T, \frac{[\text{NO}_x]}{[\text{NO}_2]}$	$[\text{OH}]$	0.769 (-10.5%)	0.219 (-13.6%)	1.013 (-5.5%)	0.324 (-6.0%)	1.129 (+0.4%)	0.375 (-0.3%)	0.853 (-8.5%)	0.251 (-9.5%)
$T, [\text{OH}]$	$\frac{[\text{NO}_x]}{[\text{NO}_2]}$	0.872 (+1.6%)	0.257 (+1.4%)	1.094 (+2.1%)	0.352 (+2.0%)	1.143 (+1.6%)	0.383 (+1.9%)	0.904 (+3.1%)	0.271 (+2.2%)

Table 2: Analysis of the effect of a vertical change in the parameters used to estimate the mean sink term in  $\text{NO}_x$  emissions: temperature, hydroxyl radical concentration, and  $\text{NO}_x:\text{NO}_2$  concentration ratio. The comparison is conducted between the estimated quantities at level  $\mathcal{B}$  and level  $\mathcal{A}$  for mask cells (MASK) and background cells (BKGD) for January, April, October and July 2019. Values within brackets represent the variation from the base case for which all quantities are estimated at level  $\mathcal{B}$ .

The transition to level  $\mathcal{A}$  generally results in a decrease in temperature, leading to an increase in the reaction rate  $k_{\text{mean}}$  and thus an increase in the emissions from the sink term. This transition has only a small influence on the total  $\text{NO}_x$  490 emission estimates, with mask and background cells emissions increasing by 4 to 6%. The influence of OH goes in the opposite direction: its concentration decreases with altitude, weakening the sink term. This weakening is particularly visible during winter months, for which the emissions are lower by up to 14%. In summer however, the effect is hardly noticeable. Finally, the influence of the  $\text{NO}_x:\text{NO}_2$  ratio is negligible on the  $\text{NO}_x$  emission estimates. Thus, the transition to level  $\mathcal{A}$  results in an increase in the sink term of 1 to 4%, due to a decrease in both concentrations 495 of NO and  $\text{NO}_2$  with respect to the vertical but with a greater decrease for  $\text{NO}_2$ . This vertical study confirms the crucial importance of the OH concentration for the accurate representation of  $\text{NO}_x$  emissions. It appears here as an important driver of the sink term, which is much more sensitive to vertical differences than temperature or the  $\text{NO}_x:\text{NO}_2$  concentration ratio.

### 4.4 Weekly cycle

500 In Egypt, the official rest day is Friday, and the economic activity of the country is *a priori* lower during this day than during the other days of the week. We therefore try to characterise this feature, by evaluating the weekly cycle of  $\text{NO}_x$  emissions. We use the TROPOMI-inferred emissions to obtain averages per day of the week. We use the quality assurance  $q_a$  of TROPOMI retrievals to ignore the days for which more than 20% of the domain has low-quality data (this happens 43 times in 2018/2019 and 28 times in 2019/2020). Such a filtering avoids accounting for the days when 505 a large part of the urban and industrial areas are covered by clouds. However, it misses situations where small clouds are localised over large emitters, in which case the corresponding emissions are under-estimated. Figure 8 shows the resulting daily emissions for the period November 2018 - November 2019 and November 2019 - November 2020.  $\text{NO}_x$  emissions are expressed as NO and in kilotons per day. A Friday minimum is observed, defining a weekly cycle. This trend is also observed for mean  $\text{NO}_2$  column densities, for which no intra-weekly variation is observed. Over the 510 2018-2019 period, Fridays have average emissions of  $0.811 \pm 0.408$  kt, which is lower than average emissions for the rest of the week, which reach  $0.997 \pm 0.533$  kt. A similar trend is observed in 2019-2020, for which the average for



515 Fridays is  $0.704 \pm 0.357$  kt and the average for other days is  $0.921 \pm 0.449$  kt. The difference in emissions between the two periods is due to smaller emissions in December 2019, January 2020 and February 2020 that are discussed in Section 4.5. On average, Friday emissions correspond to a ratio of 0.83:7 (i.e. a value of 0.83 after normalisation on the seven days of the week) for the entire domain. This result is consistent with the values obtained by Stavarakou et al. (2020), who used TROPOMI data and another emission model to calculate a ratio of 0.71:7 for Cairo and 0.89:7 for Alexandria in 2017.

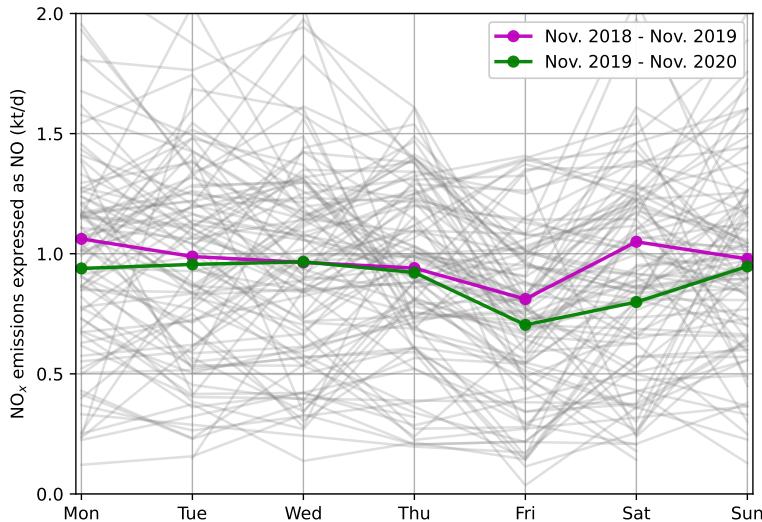


Figure 8: Weekly profiles of anthropogenic  $\text{NO}_x$  emissions for Egypt using TROPOMI observations in 2018-2019 (purple line) and 2019-2020 (green line). Thin grey lines represent individual weeks. Days for which less than 80% of the domain counts low quality observations ( $q_a < 0.75$ ) are not represented.

#### 4.5 Impacts of lockdown during COVID-19

520 The ongoing global outbreak of COVID-19 forced many countries around the world to implement unprecedented public health responses, including travel restrictions, quarantines, curfews and lockdowns. Such measures have helped to counter the spread of the virus and have, meanwhile, caused high reductions in global demand for fossil fuels (IEA, 2020). They also led to a fall in the levels of  $\text{NO}_2$  and other air pollutants across the globe (Venter et al., 2020; Bauwens et al., 2020; Gkatzelis et al., 2021). To prevent the spread of COVID-19, Egyptian authorities ordered a partial lockdown from March 15th till June 30th 2020, closing all public areas (e.g. sport centres, nightclubs, 525 restaurants and cafes) and suspending religious activities in all mosques and churches throughout the country. They also implemented more drastic measures such as a full lockdown during Easter (April 20th) and Eid (May 23rd to May 25th), before lifting some restrictions on June 1st (Hale et al., 2021). In addition to the effect of containment on the activity of the country, the global decline in consumption led to a drop in the production of certain industrial products.

530 Several studies have estimated the impact of these events on the air pollution levels in the urban centres of the country : from in-situ measurements, El-Sheekh and Hassan (2021) estimated that  $\text{NO}_2$  concentrations had dropped by 25.9% in Alexandria's city centre after the start of the lockdown on March 13th, while Abou El-Magd and Zanaty (2020) used OMI retrievals to estimate a 45.5% reduction of  $\text{NO}_2$  concentrations for the entire country during the spring compared to 2018 and 2019 average values. However, due to a changing lifetime of  $\text{NO}_2$ , reductions in the concentrations of  $\text{NO}_2$  might not be entirely due to a decrease in  $\text{NO}_x$  emissions, which leads us to focus on the 535 variation of  $\text{NO}_x$  emissions during this singular period. Using our top-down emission model, reductions in total  $\text{NO}_x$  emissions of 20.1%, 11.8% and 13.5% are observed for the months of March, April and May 2020 compared to the equivalent months in 2019. This drop of emissions in 2020 compared to 2019 calculated by the model also correspond to a decrease in observed  $\text{NO}_2$  columns: TROPOMI retrievals above mask cells show a decrease in  $\text{NO}_2$  column densities of 21.6% over the same period. However, these effects observed for the months of March, April and May 2020 are not 540 repeated in June 2020, for which emissions show an increase of 15.8% compared to June 2019. This rise is largely the result of an increase in the difference between average estimates inside and outside the mask. Indeed, emissions within the mask in June 2020 are higher than those of June 2019, due to an increase in TROPOMI concentrations

above mask cells (+7.7%) while the  $\text{NO}_2$  lifetime is almost unchanged (+3.3%). Emissions outside the mask varies in the opposite direction: a decrease in TROPOMI background concentrations (-5.4%) is observed while  $\text{NO}_2$  lifetime increases strongly (+16.0%). This increase in June emissions seems to indicate that the lift on restrictions allowed a catch-up of the economic activity which was sufficiently strong to generate higher emissions in 2020 than in 2019. Note that CAMS OH concentrations during the lockdown periods do not show significant variations from previous years, although concentration values are slightly lower in 2020 than in 2019 (about 5.5% lower over the mask cells for the period March/April/May). The near-real-time CAMS system did not take into account the decrease in anthropogenic emissions in the representation of its OH concentrations. However, the satellite constraints inherent in the system may have modulated the lockdown effects locally or globally. Given the non-linearity of the chemistry but also given the large reactivity of OH with other species whose concentrations have varied differently during the lockdown, it is difficult to determine how these observations have impacted OH concentrations.

#### 4.6 Annual cycle and comparison to inventories

Here, we attempt to compare our TROPOMI-derived  $\text{NO}_x$  emissions in Egypt to emissions from CAMS-GLOB-ANT\_v4.2 and EDGARv5.0 inventories. Figure 9 shows the total anthropogenic  $\text{NO}_x$  emissions over the mask cells from November 2018 to November 2020, calculated according to Equation (7). As indicated in Section 3.2, the emissions, calculated at 13:30 local time, are representative of the average daily consumption in Egypt. The total calculated for each month therefore corresponds to the  $\text{NO}_x$  production by human activities in the country. After aggregating the different sectors of activity, CAMS and EDGAR inventories directly provide the anthropogenic  $\text{NO}_x$  emissions over the same domain. All  $\text{NO}_x$  emissions are expressed in mass terms as NO. We note that the EDGAR inventory does not cover the period 2018-2020 (the last available year of the inventory is 2015). On Figure 9, EDGAR emissions corresponding to the period between November 2013 and November 2015 are displayed.

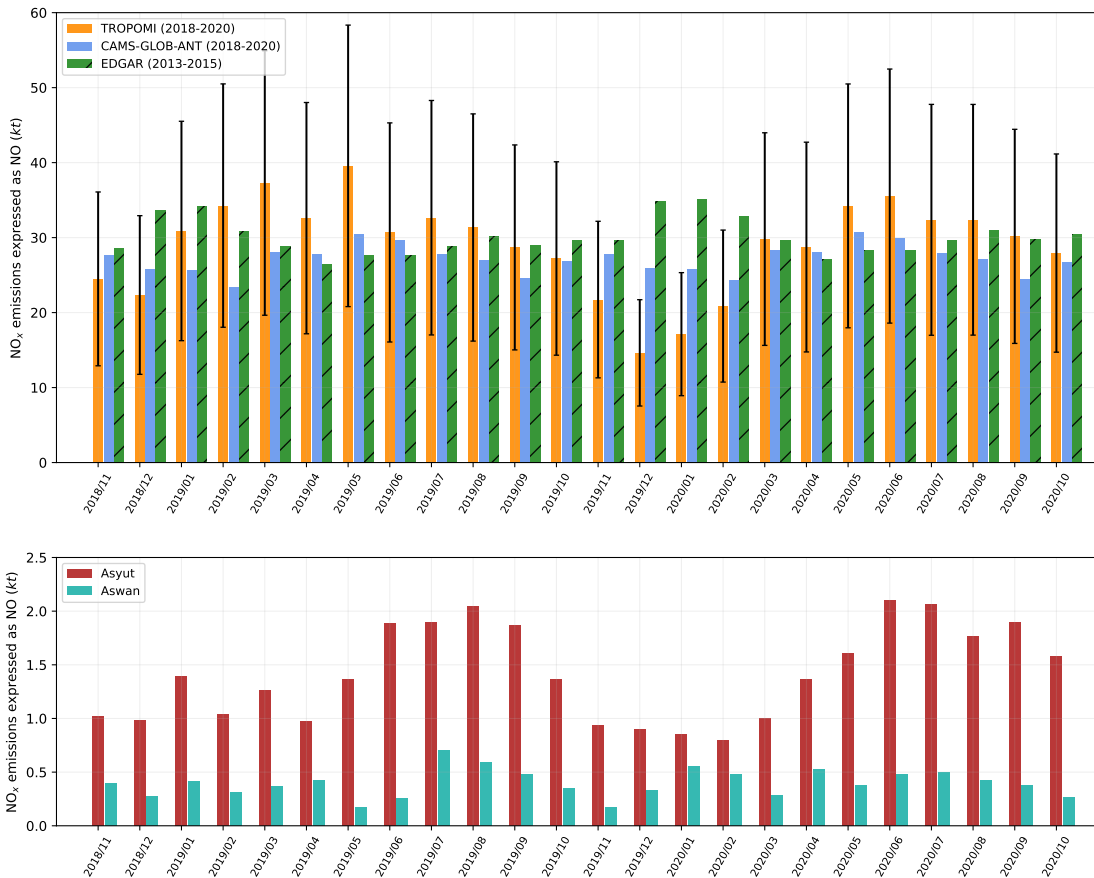


Figure 9: (top) Comparison of TROPOMI-derived anthropogenic  $\text{NO}_x$  emissions in Egypt (light blue), with the corresponding emissions from EDGAR (green with stripes) and CAMS (yellow) inventories. EDGAR data is shown for comparison purposes and covers the years 2013-2015. (bottom) TROPOMI-derived anthropogenic  $\text{NO}_x$  emissions for the cities of Asyut (dark red) and Aswan (light blue). The corresponding domains are displayed on Figure 1.



565 TROPOMI-derived emissions are higher than the CAMS inventory estimates. The top-down model estimates total  
NO<sub>x</sub> emissions of 697.6 kt over the 24 months, which is 45.9 kt higher than CAMS for the same period (651.6 kt).  
This difference is primarily localized in the first 12 months, for which TROPOMI-inferred emissions are always higher  
than the inventories and show higher values in summer than during the rest of the year. The next 12 months show  
570 similar emissions in summer but much lower values in winter. In particular, the difference is significant in December  
2019 and January 2020 (respectively 56.5% and 66.5% of CAMS levels). These emissions also contrast with other  
winter emissions, with a total of 31.7 kt for 2019-12/2020-01 against 53.3 kt for 2018-12/2019-01 and 57.7 kt for  
2020-12/2021-01. In the computations, this drop for winter 2019/2020 is mainly due to a relatively low value of  
the OH concentration which reaches  $4.61 \times 10^6$  molecules.cm<sup>-3</sup> on average for December 2019 and January 2020,  
with  $4.29 \times 10^6$  molecules.cm<sup>-3</sup> above mask cells and  $4.69 \times 10^6$  molecules.cm<sup>-3</sup> over background cells. They were  
575 respectively  $5.29$ ,  $5.74$  and  $5.18 \times 10^6$  molecules.cm<sup>-3</sup> for the previous year (2018-12/2019-01) and  $5.11$ ,  $4.90$  and  
 $5.16 \times 10^6$  molecules.cm<sup>-3</sup> for the subsequent year (2020-12/2021-01). A decrease in tropospheric columns (-18.5% for  
mask cells and -7.6% for background cells compared to winter 2018/2019) also contributes to this drop. The accuracy  
of the inferred emissions for winter 2019/2020 can therefore be questioned.

At first sight, the annual variability of TROPOMI-inferred emissions, which describes a one-year cycle with higher  
580 emissions in summer, seems to be correlated with power emissions which dominate the use of fossil fuels in Egypt  
(Abdallah and El-Shennawy, 2020). These power emissions are due to the country's residential electricity consumption  
(Attia et al., 2012; Elharidi et al., 2013; Nassief, 2014). They also meet the needs of industry. Summer peaks in  
electricity consumption are mostly driven by temperature, as illustrated by the increasing sales of air conditioning  
and ventilation systems for several decades (Wahba et al., 2018). The use of air conditioning in cars, which requires  
585 an additional amount of fuel, could also contribute to the increase of NO<sub>x</sub> emissions in summer. To support this  
hypothesis, we use our model on two smaller domains centered around the two cities of Asyut and Aswan. The  
corresponding domains are displayed on Figure 1. Both cities have similar demographic features, with populations of  
about 467,000 and 315,000 inhabitants in 2021 and human densities of about 3,000 and 1,600 inhabitants per square  
kilometer respectively. However, their industrial features largely differ. There is no large fossil fuel-fired power plant  
590 in Aswan, where most of the electricity is produced by a hydroelectric dam, whereas Asyut counts three oil and gas  
power plants of various capacities (90, 650 and 1500 MW) in its urban area. Both cities have a cement plant, but  
the one in Asyut has a larger production capacity (5.7 Mt/yr in Asyut, 0.8 Mt/yr in Aswan). Our model is used  
following the same procedure as for the main domain. The background removal is done at the scale of the country.  
A seasonal cycle appears for Asyut, with a minimum for winter months and a maximum for summer months. This  
595 cycle seems slightly shifted from the one observed for the entire country, for which May emissions are as important  
as those of summer months. We also note that the decrease in emissions for winter 2019/2020 is less marked than for  
the emissions of the whole country, and of a similar value to the previous winter. This suggests that national NO<sub>x</sub>  
emissions are indeed lower during winter, but that the values obtained for winter 2019/2020 are particularly low. We  
also find that the seasonality of the emissions is more pronounced for the Asyut domain than for the country as a  
600 whole. The case of Aswan is different. Emissions within the corresponding domain are significantly lower than for  
Asyut. The signal-to-noise ratio being higher, it is difficult to characterise an annual cycle, but the results do not seem  
to indicate low emissions in winter and high emissions in summer. This identification of a seasonal cycle identical to  
that of the entire country for a city with several power plants, and the absence of such a cycle in a city without any,  
strengthens the hypothesis that the power sector plays a major role in Egyptian NO<sub>x</sub> emissions.

605 We note that some features of the industrial activities in the country might be counteracting this trend. For some  
sectors such as cement or steel, production is lower in summer, due to the physical wear experienced by workers due  
to heat, but also due to certain periods of leave. Given the importance of industrial activities in the production of  
NO<sub>x</sub> shown in Section 4.2, this aspect cannot be neglected. The transport sector could also counteract the observed  
trend: although the use of air conditioning in cars increase NO<sub>x</sub> emissions of the sector, the observed mean traffic  
610 in the country is higher between November and February and lower between June and August, especially in Cairo  
which gathers most of the population. In the absence of additional data, it therefore seems difficult to conclude on the  
amplitude of the seasonal cycle produced by our top-down model. This caution is all the more necessary as CAMS  
and EDGAR show seasonal cycles for NO<sub>x</sub> emissions, with different dynamics than those displayed by TROPOMI  
emissions: while the EDGAR inventory predicts a maximum of emissions in December or January and a minimum  
615 in April, the CAMS inventory shows two local maxima each year in May and November and two local minima in  
February and September. The amplitude of the corresponding cycles is much lower in those inventories, representing  
14.2% of the average value for emissions estimates for EDGAR and 12.4% for CAMS. Those values must be compared  
to the amplitude displayed by TROPOMI-inferred emissions, for which the maximum/minimum ratio is about 1.8 if  
winter 2019/2020 is excluded, and 2.7 if it is included.

## 620 4.7 Uncertainties and assessments of results

The estimation of  $\text{NO}_x$  emissions is based on the use of several quantities with varying uncertainties. The error bars shown in Figures 5 and 9 are thus calculated from uncertainty statistics whose references are presented in this section. Since these references do not specify the exact nature of these statistics, we assume they correspond to standard deviations. The uncertainty of tropospheric  $\text{NO}_2$  columns under polluted conditions is dominated by the sensitivity  
625 of satellite observations to lower tropospheric air masses, expressed by the tropospheric air-mass factor (AMF). The column relative uncertainty due to the AMF is of the order of 30% (Boersma et al., 2004). S-5P validation activities indicate that TROPOMI tropospheric  $\text{NO}_2$  columns are systematically biased low by about 30%-50% over cities (Verhoelst et al., 2021), which is most likely related to the *a priori* profiles used within the operational retrieval that do not reflect well the  $\text{NO}_2$  peak close to ground. For the Middle East region, the impact of the *a priori* profile is less  
630 critical, as surface albedo is generally high and cloud fractions are generally low. Thus, we expect no such bias, and consider a relative uncertainty of 30% for the tropospheric column. Other uncertainties must be taken into account: the transition from  $\text{NO}_2$  TROPOMI columns to  $\text{NO}_x$  emissions requires parameters which appear in Equation (2) and Equation (3). For wind module, uncertainties are generally of about 1 m/s for components taken at precise altitudes (Coburn, 2019; Beirle et al., 2019). Here, we assume an uncertainty of 3 m/s for both zonal and meridional  
635 wind components. For [OH], the analysis of different methods conducted by Huijnen et al. (2019) showed smaller differences for low latitudes than for extratropics, but still significant. We thus take a relative uncertainty of 30% for OH concentration. For the reaction rate  $k_{\text{mean}}$ , the value of the corresponding relative uncertainty has been estimated by Burkholder et al. (2020). Finally, we use the sensitivity tests performed in Section 4.3 to assess the uncertainty associated with the choice of the vertical level. The cumulative effects on the final emissions of the three parameters  
640 studied, in particular the OH concentration, lead to a relative uncertainty that varies from month to month between 7 and 18%. The propagation of these different uncertainties on the monthly estimates of  $\text{NO}_x$  emissions in Egypt leads to an expanded uncertainty between 47 and 51%. For lifetimes calculated with the EMG function fitting, the corresponding expanded uncertainty ranges between 18% and 79%.

We acknowledge the fact that our treatment of  $\text{NO}_x$  is simplified. Many minor sinks highlighted in Section 3.1  
645 are not taken into account. In particular, anthropogenic VOC emissions, which remove  $\text{NO}_x$  from the atmosphere, compete with the oxidation by OH for the representation of  $\text{NO}_x$  loss. These emissions are difficult to estimate and the corresponding sink is complex to model. Taking this reaction into account would *a priori* lead to a strengthening of the sink term and thus to an increase of the  $\text{NO}_x$  emissions estimates. Moreover, due to the coarse resolution of CAMS data, OH gradients might also be underestimated, especially in the southern part of the domain, leading to  
650 a local under-estimation of the sink term and the corresponding emissions. Other assumptions in the model are also simplifications. For instance, obtaining anthropogenic emissions by subtracting the average emissions over background cells assumes that the non-anthropogenic sources of  $\text{NO}_2$  are similar inside and outside the mask, which is not true, since a large part of the mask cells correspond to croplands. For these cells, soil emissions may play a non-negligible role in the natural  $\text{NO}_2$  budget. As a consequence, mean background emissions that are removed from  $\text{NO}_x$  emissions  
655 estimates above mask cells might be under-estimated. Finally, the reliability of the data used can be questioned. The representation of the wind is crucial to avoid creating artificial patterns in the transport term. The OH concentration, which is proportional to the intensity of the sink term, is also important. We have shown that OH concentrations are partially responsible for an important drop in  $\text{NO}_x$  emissions in the winter of 2019/2020 that may be unrealistic. Because this decrease is largely due to variations in OH concentrations provided by CAMS, whose reliability has  
660 been evaluated for Riyadh, then the transposability hypothesis between Riyadh and Egypt may be subject to further discussion.

## 5 Conclusions

In this study, we investigated the potential of a top-down model of  $\text{NO}_x$  emissions based on TROPOMI retrievals at high resolution over Egypt. The model is based on the study of a transport term and a sink term that requires  
665 different parameters to be calculated. Among those parameters, the concentration in OH, involved in the calculation of the  $\text{NO}_2$  mixed lifetime, is of fundamental importance. The comparison between  $\text{NO}_2$  lifetimes derived from OH concentrations and  $\text{NO}_2$  lifetimes derived from EMG function fittings of line density profiles shows that the OH concentration provided by CAMS is reasonably reliable for the country. Parameters are taken in the first 200 m of the planetary boundary layer, because it is where OH shows the best consistency. However, the vertical sensitivity  
670 linked to this parameters remains high. Results illustrate the importance of the transport term at local scale, which is of the same order of magnitude as the sink term above large cities and industrial facilities; it ceases to be relevant only at the scale of the whole country. The top-down model is able to characterise declines in human activities due

to restrictions during the COVID-19 pandemic or to Friday rest. It also estimates higher emissions during summer. These high emissions may be interpreted by a higher consumption of electricity driven by air-conditioning during hot days, but it remains unclear whether this pattern clearly reproduces changes in human activity, in particular because the emission inventories show different seasonalities. These inventories also differ in the amount of total emissions: the average value for TROPOMI-derived NO<sub>x</sub> emissions is 7.0% higher than CAMS-GLOB-ANT\_v4.2 estimates. This discrepancy could be resolved by comparing the results of the model and inventory estimates to industrial production or electricity consumption data at the scale of countries or regions.

Here, our estimation of NO<sub>x</sub> emissions benefited from favorable conditions. Egypt has a desertic climate, allowing to neglect many NO<sub>x</sub> loss mechanisms for the sink term calculation, a flat terrain on most of its territory, limiting wind field errors for the transport term calculation, and a large population concentrated in a small number of cities, providing NO<sub>2</sub> maps with large signal-to-noise ratios above urban and industrial areas. For other regions of the world that do not have such features, the method presented here must be modified accordingly. However, we expect this method to be applicable to most countries similar to Egypt without substantial changes. For Middle East countries, this study thus demonstrates the potential of TROPOMI data for evaluating NO<sub>x</sub> emissions. More generally, it demonstrates the importance of the contribution of independent observation systems to overcome the weaknesses of emission inventories, provided that the local chemistry is well understood and modelled. The development of similar applications for different species is likely to allow a better monitoring of global anthropogenic emissions, therefore helping companies and countries to report their emissions of air pollutants and greenhouse gases as part of their strategies and obligations to tackle air pollution issues and climate change.

#### Data availability.

TROPOMI product: <http://www.tropomi.eu/data-products/data-access>

CAMS NRT: <https://ads.atmosphere.copernicus.eu/cdsapp!/dataset/cams-global-atmospheric-composition-forecasts>

ERA5 reanalysis: <https://cds.climate.copernicus.eu/cdsapp!/dataset/reanalysis-era5-pressure-levels-monthly-means>

Global Rural-Urban Mapping Project (GRUMP): <https://sedac.ciesin.columbia.edu/data/collection/grump-v1>

Oil and gas power plants: <http://globalenergyobservatory.org/>

Industrial facilities: <https://www.industryabout.com>

Flaring sites: [https://eogdata.mines.edu/download\\_global\\_flare.html](https://eogdata.mines.edu/download_global_flare.html)

CAMS-GLOB-ANT\_v4.2: <https://permalink.aeris-data.fr/CAMS-GLOB-ANT>

EDGARv5.0: [https://edgar.jrc.ec.europa.eu/dataset\\_ap50](https://edgar.jrc.ec.europa.eu/dataset_ap50)

**Competing interests.** The authors declare that they have no conflict of interest.

**Acknowledgements.** The authors would like to thank Steven J. Davis (University of California, Irvine) and Dan Tong (Tsinghua University) for their contribution to the construction of our emitters database.

**Financial support.** This study has been funded by the European Union's Horizon 2020 research and innovation programme under grant agreement N° 856612 (EMME-CARE).

## References

Abdallah, L. and El-Shennawy, T. (2020). Evaluation of CO<sub>2</sub> emission from Egypt's future power plants. *Euro-Mediterranean Journal for Environmental Integration*, 5(3):1–8.

About El-Magd, I. and Zanaty, N. (2020). Impacts of short-term lockdown during COVID-19 on air quality in Egypt. *The Egyptian Journal of Remote Sensing and Space Science*.

Attia, S., Evrard, A., and Gratia, E. (2012). Development of benchmark models for the Egyptian residential buildings sector. *Applied Energy*, 94:270–284.

Baklanov, A., Molina, L. T., and Gauss, M. (2016). Megacities, air quality and climate. *Atmospheric Environment*, 126:235–249.

- Bauwens, M., Compernolle, S., Stavrakou, T., Müller, J.-F., Van Gent, J., Eskes, H., Levelt, P. F., van der A, R., Veeffkind, J., Vlietinck, J., et al. (2020). Impact of coronavirus outbreak on NO<sub>2</sub> pollution assessed using TROPOMI and OMI observations. *Geophysical Research Letters*, 47(11):e2020GL087978.
- 720 Beirle, S., Boersma, K. F., Platt, U., Lawrence, M. G., and Wagner, T. (2011). Megacity emissions and lifetimes of nitrogen oxides probed from space. *Science*, 333(6050):1737–1739.
- Beirle, S., Borger, C., Dörner, S., Eskes, H., Kumar, V., de Laat, A., and Wagner, T. (2021). Catalog of NO<sub>x</sub> emissions from point sources as derived from the divergence of the NO<sub>2</sub> flux for TROPOMI. *Earth System Science Data*, 13(6):2995–3012.
- 725 Beirle, S., Borger, C., Dörner, S., Li, A., Hu, Z., Liu, F., Wang, Y., and Wagner, T. (2019). Pinpointing nitrogen oxide emissions from space. *Science advances*, 5(11):eaax9800.
- Boersma, K., Eskes, H., and Brinksma, E. (2004). Error analysis for tropospheric NO<sub>2</sub> retrieval from space. *Journal of Geophysical Research: Atmospheres*, 109(D4).
- Burkholder, J., Sander, S., Abbatt, J., Barker, J., Cappa, C., Crouse, J., Dibble, T., Huie, R., Kolb, C., Kurylo, M., et al. (2020). Chemical kinetics and photochemical data for use in atmospheric studies; evaluation number 19. Technical report, Pasadena, CA: Jet Propulsion Laboratory, National Aeronautics and Space . . . .
- 730 Butkovskaya, N., Kukui, A., Pouvesle, N., and Le Bras, G. (2005). Formation of nitric acid in the gas-phase HO<sub>2</sub>+NO reaction: Effects of temperature and water vapor. *The Journal of Physical Chemistry A*, 109(29):6509–6520.
- Butkovskaya, N., Rayez, M.-T., Rayez, J.-C., Kukui, A., and Le Bras, G. (2009). Water vapor effect on the HNO<sub>3</sub> yield in the HO<sub>2</sub>+NO reaction: experimental and theoretical evidence. *The Journal of Physical Chemistry A*, 113(42):11327–11342.
- 735 Coburn, J. J. (2019). Assessing wind data from reanalyses for the upper midwest. *Journal of Applied Meteorology and Climatology*, 58(3):429–446.
- Crippa, M., Guizzardi, D., Muntean, M., Schaaf, E., Solazzo, E., Monforti-Ferrario, F., Olivier, J., and Vignati, E. (2020). Fossil CO<sub>2</sub> emissions of all world countries. *Luxembourg: European Commission*, pages 1–244.
- 740 Crippa, M., Janssens-Maenhout, G., Dentener, F., Guizzardi, D., Sindelarova, K., Muntean, M., Van Dingenen, R., and Granier, C. (2016). Forty years of improvements in European air quality: regional policy-industry interactions with global impacts. *Atmospheric Chemistry and Physics*, 16(6):3825–3841.
- de Foy, B., Lu, Z., Streets, D. G., Lamsal, L. N., and Duncan, B. N. (2015). Estimates of power plant NO<sub>x</sub> emissions and lifetimes from OMI NO<sub>2</sub> satellite retrievals. *Atmospheric Environment*, 116:1–11.
- 745 Delaria, E. R., Place, B. K., Liu, A. X., and Cohen, R. C. (2020). Laboratory measurements of stomatal NO<sub>2</sub> deposition to native california trees and the role of forests in the NO<sub>x</sub> cycle. *Atmospheric Chemistry and Physics*, 20(22):14023–14041.
- Ehhalt, D. H., Rohrer, F., and Wahner, A. (1992). Sources and distribution of NO<sub>x</sub> in the upper troposphere at northern mid-latitudes. *Journal of Geophysical Research: Atmospheres*, 97(D4):3725–3738.
- 750 El-Magd, A., Zanaty, N., Ali, E. M., Irie, H., Abdelkader, A. I., et al. (2020). Investigation of aerosol climatology, optical characteristics and variability over Egypt based on satellite observations and in-situ measurements. *Atmosphere*, 11(7):714.
- El-Sheekh, M. M. and Hassan, I. A. (2021). Lockdowns and reduction of economic activities during the COVID-19 pandemic improved air quality in Alexandria, Egypt. *Environmental Monitoring and Assessment*, 193(1):1–7.
- 755 Elharidi, A. M. A. H., Tuohy, P. G., and Teamah, M. (2013). Facing the growing problem of the electric power consumption in Egyptian residential building using building performance simulation program. In *Building simulation Cairo 2013 conference*.
- Elvidge, C. D., Zhizhin, M., Baugh, K., Hsu, F.-C., and Ghosh, T. (2016). Methods for global survey of natural gas flaring from visible infrared imaging radiometer suite data. *Energies*, 9(1):14.

- 760 EPA, U. (2016). Integrated science assessment for oxides of nitrogen–health criteria. *US Environmental Protection Agency, Washington, DC [Google Scholar]*.
- Eskes, H., Eichmann, K., Lambert, J., Loyola, D., Veeffkind, J., Dehn, A., and Zehner, C. (2019). S5P Mission Performance Centre Nitrogen Dioxide [L2\_NO2] readme. *Royal Netherlands Meteorological Institute (KNMI) De Bilt, the Netherlands, version, 1(00)*.
- 765 Fang, H., Baret, F., Plummer, S., and Schaepman-Strub, G. (2019). An overview of global leaf area index (LAI): Methods, products, validation, and applications. *Reviews of Geophysics*, 57(3):739–799.
- Filioglou, M., Giannakaki, E., Backman, J., Kesti, J., Hirsikko, A., Engelmann, R., O’Connor, E., Leskinen, J. T., Shang, X., Korhonen, H., et al. (2020). Optical and geometrical aerosol particle properties over the United Arab Emirates. *Atmospheric Chemistry and Physics*, 20(14):8909–8922.
- 770 Fischer, E., Jacob, D. J., Yantosca, R. M., Sulprizio, M. P., Millet, D., Mao, J., Paulot, F., Singh, H., Roiger, A., Ries, L., et al. (2014). Atmospheric peroxyacetyl nitrate (PAN): a global budget and source attribution. *Atmospheric Chemistry and Physics*, 14(5):2679–2698.
- Friedrich, N., Eger, P., Shenolikar, J., Sobanski, N., Schuladen, J., Dienhart, D., Hottmann, B., Tadic, I., Fischer, H., Martinez, M., et al. (2021). Reactive nitrogen around the arabian peninsula and in the mediterranean sea during  
775 the 2017 AQABA ship campaign. *Atmospheric Chemistry and Physics*, 21(10):7473–7498.
- Gkatzelis, G. I., Gilman, J. B., Brown, S. S., Eskes, H., Gomes, A. R., Lange, A. C., McDonald, B. C., Peischl, J., Petzold, A., Thompson, C. R., et al. (2021). The global impacts of COVID-19 lockdowns on urban air pollution: a critical review and recommendations. *Elementa: Science of the Anthropocene*, 9(1).
- Goldberg, D. L., Lu, Z., Streets, D. G., de Foy, B., Griffin, D., McLinden, C. A., Lamsal, L. N., Krotkov, N. A., and  
780 Eskes, H. (2019). Enhanced Capabilities of TROPOMI NO<sub>2</sub>: Estimating NO<sub>x</sub> from North American Cities and Power Plants. *Environmental science & technology*, 53(21):12594–12601.
- Graedel, T., Farrow, L., and Weber, T. (1976). Kinetic studies of the photochemistry of the urban troposphere. *Atmospheric Environment (1967)*, 10(12):1095–1116.
- Granier, C., Darras, S., van der Gon, H. D., Jana, D., Elguindi, N., Bo, G., Michael, G., Marc, G., Jalkanen, J.-P.,  
785 Kuenen, J., et al. (2019). *The Copernicus atmosphere monitoring service global and regional emissions (April 2019 version)*. PhD thesis, Copernicus Atmosphere Monitoring Service.
- Guenther, A., Karl, T., Harley, P., Wiedinmyer, C., Palmer, P. I., and Geron, C. (2006). Estimates of global terrestrial isoprene emissions using MEGAN (model of emissions of gases and aerosols from nature). *Atmospheric Chemistry and Physics*, 6(11):3181–3210.
- 790 Hale, T., Angrist, N., Goldszmidt, R., Kira, B., Petherick, A., Phillips, T., Webster, S., Cameron-Blake, E., Hallas, L., Majumdar, S., et al. (2021). A global panel database of pandemic policies (oxford covid-19 government response tracker). *Nature Human Behaviour*, 5(4):529–538.
- Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas, J., Peubey, C., Radu, R., Schepers, D., et al. (2020). The ERA5 global reanalysis. *Quarterly Journal of the Royal Meteorological Society*,  
795 146(730):1999–2049.
- Hoelzemann, J. J., Schultz, M. G., Brasseur, G. P., Granier, C., and Simon, M. (2004). Global wildland fire emission model (GWEM): Evaluating the use of global area burnt satellite data. *Journal of Geophysical Research: Atmospheres*, 109(D14).
- Hoesly, R. M., Smith, S. J., Feng, L., Klimont, Z., Janssens-Maenhout, G., Pitkanen, T., Seibert, J. J., Vu, L., Andres, R. J., Bolt, R. M., et al. (2018). Historical (1750–2014) anthropogenic emissions of reactive gases and aerosols from the community emissions data system (CEDS). *Geoscientific Model Development*, 11(1):369–408.  
800
- Huijnen, V., Eskes, H., Wagner, A., Schulz, M., Christophe, Y., Ramonet, M., Basart, S., Benedictow, A., Blechschmidt, A.-M., Chabrillat, S., et al. (2016). Validation report of the CAMS near-real-time global atmospheric composition service: System evolution and performance statistics. Status up to 1 June 2016;  
805 [https://pure.mpg.de/rest/items/item\\_2441827/component/file\\_2441834/content](https://pure.mpg.de/rest/items/item_2441827/component/file_2441834/content).

- Huijnen, V., Pozzer, A., Arteta, J., Brasseur, G., Bouarar, I., Chabrillat, S., Christophe, Y., Doumbia, T., Fleming, J., Guth, J., et al. (2019). Quantifying uncertainties due to chemistry modelling—evaluation of tropospheric composition simulations in the CAMS model (cycle 43r1). *Geoscientific Model Development*, 12(4):1725–1752.
- IEA, U. (2020). Global energy review 2020. *Ukraine*. [Online] <https://www.iea.org/countries/ukraine> [Accessed: 2020-09-10].
- 810
- Jaeglé, L., Steinberger, L., Martin, R. V., and Chance, K. (2005). Global partitioning of NO<sub>x</sub> sources using satellite observations: Relative roles of fossil fuel combustion, biomass burning and soil emissions. *Faraday discussions*, 130:407–423.
- Janssens-Maenhout, G., Crippa, M., Guizzardi, D., Muntean, M., Schaaf, E., Dentener, F., Bergamaschi, P., Pagliari, V., Olivier, J. G., Peters, J. A., et al. (2019). Edgar v4. 3.2 global atlas of the three major greenhouse gas emissions for the period 1970–2012. *Earth System Science Data*, 11(3):959–1002.
- 815
- Khoder, M. (2009). Diurnal, seasonal and weekdays–weekends variations of ground level ozone concentrations in an urban area in greater cairo. *Environmental Monitoring and Assessment*, 149(1):349–362.
- Kim, H. C., Bae, C., Bae, M., Kim, O., Kim, B.-U., Yoo, C., Park, J., Choi, J., Lee, J.-b., Lefer, B., et al. (2020). Space-Borne monitoring of NO<sub>x</sub> emissions from cement kilns in South Korea. *Atmosphere*, 11(8):881.
- 820
- Lama, S., Houweling, S., Boersma, K. F., Eskes, H., Aben, I., Denier van der Gon, H. A., Krol, M. C., Dolman, H., Borsdorff, T., and Lorente, A. (2020). Quantifying burning efficiency in megacities using the NO<sub>2</sub>/CO ratio from the Tropospheric Monitoring Instrument (TROPOMI). *Atmospheric Chemistry and Physics*, 20(17):10295–10310.
- Lange, K., Richter, A., and Burrows, J. P. (2021). Variability of nitrogen oxide emission fluxes and lifetimes estimated from Sentinel-5P TROPOMI observations. *Atmospheric Chemistry and Physics Discussions*, pages 1–32.
- 825
- Lelieveld, J., Gromov, S., Pozzer, A., and Taraborrelli, D. (2016). Global tropospheric hydroxyl distribution, budget and reactivity. *Atmospheric Chemistry and Physics*, 16(19):12477–12493.
- Leue, C., Wenig, M., Wagner, T., Klimm, O., Platt, U., and Jähne, B. (2001). Quantitative analysis of NO<sub>x</sub> emissions from global ozone Monitoring Experiment satellite image sequences. *Journal of Geophysical Research: Atmospheres*, 106(D6):5493–5505.
- 830
- Levy, H. (1971). Normal atmosphere: Large radical and formaldehyde concentrations predicted. *Science*, 173(3992):141–143.
- Li, M., Karu, E., Brenninkmeijer, C., Fischer, H., Lelieveld, J., and Williams, J. (2018). Tropospheric OH and stratospheric OH and Cl concentrations determined from CH<sub>4</sub>, CH<sub>3</sub>Cl, and SF<sub>6</sub> measurements. *NPJ Climate and Atmospheric Science*, 1(1):1–7.
- 835
- Lin, J.-T. (2012). Satellite constraint for emissions of nitrogen oxides from anthropogenic, lightning and soil sources over East China on a high-resolution grid. *Atmospheric Chemistry and Physics*, 12(6):2881–2898.
- Logan, J. A., Prather, M. J., Wofsy, S. C., and McElroy, M. B. (1981). Tropospheric chemistry: a global perspective. *Journal of Geophysical Research: Oceans*, 86(C8):7210–7254.
- 840
- Longfellow, C. A., Ravishankara, A., and Hanson, D. R. (1999). Reactive uptake on hydrocarbon soot: Focus on NO<sub>2</sub>. *Journal of Geophysical Research: Atmospheres*, 104(D11):13833–13840.
- Lorente, A., Boersma, K., Eskes, H., Veefkind, J., Van Geffen, J., De Zeeuw, M., van der Gon, H. D., Beirle, S., and Krol, M. (2019). Quantification of nitrogen oxides emissions from build-up of pollution over Paris with TROPOMI. *Scientific reports*, 9(1):1–10.
- 845
- Lorente, A., Folkert Boersma, K., Yu, H., Dörner, S., Hilboll, A., Richter, A., Liu, M., Lamsal, L. N., Barkley, M., Smedt, I. D., et al. (2017). Structural uncertainty in air mass factor calculation for NO<sub>2</sub> and HCHO satellite retrievals. *Atmospheric Measurement Techniques*, 10(3):759–782.
- Martin, R. V., Jacob, D. J., Chance, K., Kurosu, T. P., Palmer, P. I., and Evans, M. J. (2003). Global inventory of nitrogen oxide emissions constrained by space-based observations of NO<sub>2</sub> columns. *Journal of Geophysical Research: Atmospheres*, 108(D17).
- 850

- Mijling, B. and Van Der A, R. (2012). Using daily satellite observations to estimate emissions of short-lived air pollutants on a mesoscopic scale. *Journal of Geophysical Research: Atmospheres*, 117(D17).
- Moxim, W., Levy, H., and Kasibhatla, P. (1996). Simulated global tropospheric PAN: Its transport and impact on NOx. *Journal of Geophysical Research: Atmospheres*, 101(D7):12621–12638.
- 855 Müller, J.-F. and Stavrakou, T. (2005). Inversion of CO and NOx emissions using the adjoint of the images model. *Atmospheric Chemistry and Physics*, 5(5):1157–1186.
- Nassief, M. M. (2014). Evaluation of electricity consumption of a residential flat in egypt. *American Journal of Electrical Power and Energy Systems*, 3(2):7–44.
- Nault, B. A., Garland, C., Wooldridge, P. J., Brune, W. H., Campuzano-Jost, P., Crounse, J. D., Day, D. A., Dibb, J., 860 Hall, S. R., Huey, L. G., et al. (2016). Observational constraints on the oxidation of NOx in the upper troposphere. *The Journal of Physical Chemistry A*, 120(9):1468–1478.
- Rohrer, F. and Berresheim, H. (2006). Strong correlation between levels of tropospheric hydroxyl radicals and solar ultraviolet radiation. *Nature*, 442(7099):184–187.
- Romer Present, P. S., Zare, A., and Cohen, R. C. (2020). The changing role of organic nitrates in the removal and 865 transport of NOx. *Atmospheric Chemistry and Physics*, 20(1):267–279.
- Sander, S., Friedl, R., Abbatt, J., Barker, J., Burkholder, J., Golden, D., Kolb, C., Kurylo, M., Moortgat, G., Wine, P., et al. (2011). Chemical kinetics and photochemical data for use in atmospheric studies, jpl publication 10-6. *Evaluation*, (17).
- Seinfeld, J. H. (1989). Urban air pollution: state of the science. *Science*, 243(4892):745–752.
- 870 Seinfeld, J. H. and Pandis, S. N. (2006). Atmospheric chemistry and physics from air pollution to climate change.
- Singh, A. and Agrawal, M. (2007). Acid rain and its ecological consequences. *Journal of Environmental Biology*, 29(1):15.
- Sobanski, N., Thieser, J., Schuladen, J., Sauvage, C., Song, W., Williams, J., Lelieveld, J., and Crowley, J. N. (2017). Day and night-time formation of organic nitrates at a forested mountain site in south-west germany. *Atmospheric 875 Chemistry and Physics*, 17(6):4115–4130.
- Stavrakou, T., Müller, J.-F., Bauwens, M., Boersma, K., and van Geffen, J. (2020). Satellite evidence for changes in the NO2 weekly cycle over large cities. *Scientific reports*, 10(1):1–9.
- Stavrakou, T., Müller, J.-F., Boersma, K., Van Der A, R., Kurokawa, J., Ohara, T., and Zhang, Q. (2013). Key chemical NOx sink uncertainties and how they influence top-down emissions of nitrogen oxides. *Atmospheric Chemistry 880 and Physics*, 13(17):9057–9082.
- Tang, L., Xue, X., Jia, M., Jing, H., Wang, T., Zhen, R., Huang, M., Tian, J., Guo, J., Li, L., et al. (2020). Iron and steel industry emissions and contribution to the air quality in China. *Atmospheric Environment*, 237:117668.
- CIESIN (2019). CIESIN. Global Rural-Urban Mapping Project, Version 1 (GRUMPv1). Center for International Earth Science Information Network - CIESIN - Columbia University, International Food Policy Research Institute - IFPRI, The World Bank, and Centro Internacional de Agricultura Tropical - CIAT. 2011, Palisades, NY: NASA Socioeconomic Data and Applications Center (SEDAC).
- 885
- EEHC (2021). Egyptian Electricity Holding Company annual report 2019/2020.
- UNEP (2015). Air quality policies in Egypt.
- Valin, L., Russell, A., and Cohen, R. (2013). Variations of OH radical in an urban plume inferred from NO2 column 890 measurements. *Geophysical Research Letters*, 40(9):1856–1860.
- Valin, L., Russell, A., Hudman, R., and Cohen, R. (2011). Effects of model resolution on the interpretation of satellite NO2 observations. *Atmospheric Chemistry and Physics*, 11(22):11647–11655.

- 895 Veefkind, J., Aben, I., McMullan, K., Förster, H., De Vries, J., Otter, G., Claas, J., Eskes, H., De Haan, J., Kleipool, Q., et al. (2012). TROPOMI on the ESA Sentinel-5 Precursor: A GMES mission for global observations of the atmospheric composition for climate, air quality and ozone layer applications. *Remote sensing of environment*, 120:70–83.
- Venter, Z. S., Aunan, K., Chowdhury, S., and Lelieveld, J. (2020). COVID-19 lockdowns cause global air pollution declines. *Proceedings of the National Academy of Sciences*, 117(32):18984–18990.
- 900 Verhoelst, T., Compornolle, S., Pinardi, G., Lambert, J.-C., Eskes, H. J., Eichmann, K.-U., Fjæraa, A. M., Granville, J., Niemeijer, S., Cede, A., et al. (2021). Ground-based validation of the copernicus sentinel-5p TROPOMI NO<sub>2</sub> measurements with the NDACC ZSL-DOAS, MAX-DOAS and pandonia global networks. *Atmospheric Measurement Techniques*, 14(1):481–510.
- Wahba, S. M., Kamel, B. A., Nassar, K. M., and Abdelsalam, A. S. (2018). Effectiveness of green roofs and green walls on energy consumption and indoor comfort in arid climates. *Civil Engineering Journal*, 4(10):2284–2295.
- 905 Wiedinmyer, C., Tie, X., Guenther, A., Neilson, R., and Granier, C. (2006). Future changes in biogenic isoprene emissions: how might they affect regional and global atmospheric chemistry? *Earth Interactions*, 10(3):1–19.
- Wolfe, G. M., Nicely, J. M., Clair, J. M. S., Hanisco, T. F., Liao, J., Oman, L. D., Brune, W. B., Miller, D., Thames, A., Abad, G. G., et al. (2019). Mapping hydroxyl variability throughout the global remote troposphere via synthesis of airborne and satellite formaldehyde observations. *Proceedings of the National Academy of Sciences*, 116(23):11171–11180.
- 910 Xue, B. and Ren, W. (2012). China’s uncertain CO<sub>2</sub> emissions. *Nature Climate Change*, 2(11):762–762.
- Yienger, J. and Levy, H. (1995). Empirical model of global soil-biogenic NO<sub>x</sub> emissions. *Journal of Geophysical Research: Atmospheres*, 100(D6):11447–11464.