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Quantifying NO$_\mathrm{x}$ emissions in Egypt using TROPOMI observations

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Abstract. Urban areas and industrial facilities, which concentrate the majority of human activity and industrial production, are major sources of air pollutants, with serious implications for human health and global climate. For most of these pollutants, emission inventories are often highly uncertain, especially in developing countries. Spaceborne measurements from the TROPOMI instrument, onboard the Sentinel-5 Precursor satellite, are used to retrieve nitrogen dioxide (NO$_2$) column densities at high spatial resolution. Here, we use two years of TROPOMI retrievals to map nitrogen oxides (NO$_\mathrm{x}$ = NO + NO$_2$) emissions in Egypt with a top-down approach using the continuity equation in steady state. Emissions are expressed as the sum of a transport term and a sink term representing the three-body reaction comprising NO$_2$ and hydroxyl radical (OH). This sink term requires information on the lifetime of NO$_2$, which is calculated with the use of the CAMS near-real-time temperature and OH concentration fields. We compare this derived lifetime with the lifetime inferred from the fitting of NO$_2$ line density profiles in large plumes with an exponentially modified Gaussian function. This comparison, which is conducted for different samples of NO$_2$ patterns above the city of Riyadh, provides information on the reliability of the CAMS near-real-time OH concentration fields; it also provides the location of the most appropriate vertical level to represent typical pollution sources in industrial areas and megacities in the Middle East region. In Egypt, total emissions of NO$_\mathrm{x}$ are dominated by the sink term, but they can be locally dominated by wind transport, especially along the Nile where human activities are concentrated. Megacities and industrial regions clearly appear as the largest sources of NO$_\mathrm{x}$ emissions in the country. Our top-down model infers emissions with a marked annual variability. By looking at the spatial distribution of emissions at the scale of different cities with different industrial characteristics, it appears that this variability is consistent with national electricity consumption. We detect lower emissions on Fridays, which are inherent to the social norm of the country, and quantify the drop in emissions in 2020 due to the COVID-19 pandemic. Overall, our estimations of NO$_\mathrm{x}$ emissions for Egypt are 7.0% higher than the CAMS-GLOB-ANT_v4.2 inventory, and significantly differ in terms of seasonality.

1 Introduction

Economic growth in developing countries has led to a strong increase of urban air pollution (Baklanov et al., 2016 [1]). Among the different pollutants, nitrogen oxides are key species. They are generally the products of fuel combustion, such as the burning of hydrocarbons in the air at high temperature. The main sources of these compounds are vehicle engines, but also heavy industrial facilities such as power plants, iron and steel mills (Tang et al., 2020 [2]) and cement kilns (Kim et al., 2020 [3]). Their accumulation in the lowest layers of the troposphere contributes to the formation of smog and acid rain (Singh et al., 2007 [4]). They also have a significant effect on human health, as they can cause various respiratory diseases (EPA, US., 2016 [5]). To deal with these phenomena, national and regional governments generally enact a series of air pollution control strategies, which typically take the form of bans on certain polluting technologies, with the aim of reducing the concentration of pollutants at the local level to targets that must be achieved within a given timeframe. These strategies, which also help driving technological innovation, have had a significant effect in Europe (Crippa et al., 2016 [6]).

In Egypt, population growth, urbanisation, socio-economic development and the associated increase in the vehicle

fleet led to a major degradation of air quality in the last decades, especially in highly populated areas such as Greater Cairo and the Nile Delta (Abou El-Magd et al., 2020 [7]) which gather the majority of the population. The Ministry of State for the Environment has thus initiated new policies that aim to reduce pollution levels throughout the country, through technical mitigation of emissions, emission standards for vehicles and intersectoral collaboration (UNEP, 2015 [8]). However, Egypt, like most developing countries, lacks the local infrastructure to access detailed information on technical factors such as energy consumption or fuel type and technology, leading to discrepancies in inventories (Xue et al., 2012 [9]). As a consequence, the monitoring of emissions, which is important to evaluate the effects of air pollution control policies, is of limited reliability.

To overcome the uncertainties in the emission inventories, the use of independent observations systems is becoming increasingly prevalent. In this study, we investigate the use of satellite remote sensing of atmospheric concentrations to improve the quantification of NO_x emissions in Egypt. Spectrally resolved satellite measurements of solar backscattered radiation enable the quantification of NO_2 and other trace gases absorbing in the UV-Visible spectral range based on their characteristic spectral absorption patterns. Tropospheric vertical column densities, i.e. vertically integrated NO_2 concentrations in the troposphere, have been providing information on the spatial distribution of tropospheric NO_2 at global scale for nearly 30 years, allowing the identification of different sources of NO_x and the quantification of the associated emissions (Leue et al., 2001 [10]; Martin et al., 2003 [11]; Mijling and van der A, 2012 [12]; de Foy et al., 2015 [13]; Goldberg et al., 2019 [14]; Beirle et al., 2019 [15]; Lorente et al., 2019 [16]; Lange et al., 2021 [17]). In October 2017, the Sentinel-5 Precursor satellite was launched. Its main instrument is the TROPOspheric Monitoring Instrument (TROPOMI), which provides tropospheric NO_2 column densities at high spatial resolution with a large swath width and with a daily frequency (Veefkind et al., 2012 [18]). By applying the steady-state continuity equation (Beirle et al., 2019 [15]; Lama et al., [19], 2020), it is possible to build a top-down model that directly quantifies NO_x emissions from these NO_2 column densities, provided that some key parameters (wind, temperature, hydroxyl radical concentration and concentration ratio between NO_x and NO_2) are correctly estimated. This model is used to quantify the anthropogenic NO_x emissions in Egypt for a 2-year period, from November 2018 to November 2020.

This paper is organised as follows: Section 2 provides a description of the datasets used in this study. Section 3 explains the build-up and the limits of the top-down approach used to quantify NO_x emissions in Egypt. It also presents a method for validating the model parameters by using NO_2 line density profiles over Riyadh, Saudi Arabia. Section 4 presents the analysis of this validation method. It presents the location of the main NO_x sources in Egypt and evaluates the vertical sensitivity of the model. It also assesses the ability of the model to show less human activity on Fridays and during the lockdown that took place during the COVID-19 pandemic. It finally confronts the inferred emissions with different inventories in terms of amplitude and seasonality. Section 5 presents our conclusion and general remarks.

2 Instrumentation and data

2.1 TROPOMI NO_2 retrievals

The TROPOspheric Atmosphere Monitoring Instrument (TROPOMI), onboard the European Space Agency's (ESA) Sentinel-5 Precursor (S-5P) satellite, provides measurements for atmospheric composition. TROPOMI is a spectrometer observing wavelengths in the infrared, visible and ultraviolet light at around 13:30 local time. The UV-Visible spectral band at 405-465 nm is used to retrieve NO_2. Other spectral bands are used to observe methane, formaldehyde, sulphur dioxide, carbon monoxide and ozone, as well as aerosols and cloud physical properties. The very high spatial resolution offered by TROPOMI (originally 3.5×7 km^2 at nadir, improved to 3.5×5.5 km^2 since 6 August 2019) provides unprecedented information on tropospheric NO_2. Its large swath width (~ 2600 km) makes it possible to construct NO_2 images on large spatial scales. Those images greatly improve the potential for detecting highly localised pollution plumes above the ground, identifying small-scale emission sources but also estimating emissions from megacities, industrial facilities and biomass burning. We use TROPOMI NO_2 retrievals (Level 2 data, OFFL stream) from November 2018 to November 2020 over Egypt. We also use them over Saudi Arabia, and more specifically over the city of Riyadh, to evaluate the reliability of other parameters. This will be explained in Section 3.3. Both countries have an arid climate, which offers a large number of clear-sky days throughout the year, enabling the calculation of monthly averages based on more than 20 observations. They are also the host to many large plumes of pollutants due to high human concentrations along rivers and around megacities, which allows us to observe high NO_2 concentration patterns with a high signal-to-noise ratio. TROPOMI products provide a quality assurance value q_a, which ranges from 0 (no data) to 1 (high-quality data). For our analysis of concentrations, we selected NO_2 retrievals with q_a values greater than 0.75, which systematically correspond to clear-sky conditions (Eskes et al., 2019 [20]).

TROPOMI soundings are gridded at a spatial resolution of $0.1° \times 0.1°$ with daily coverage. This resolution is lower than that of the instrument; the gridding thus provides a grid for which most NO_2 columns correspond to one or more measurements. The observed plumes remain correctly resolved. Cells without measurements are infrequent, which facilitates the calculation of derivatives.

2.2 Wind data

The horizontal wind $\mathbf{w} = (u, v)$ is taken from the European Centre for Medium-Range Weather Forecasts (ECMWF) ERA5 data archive (fifth generation of atmospheric reanalyses) at a horizontal resolution of $0.25° \times 0.25°$ on 37 pressure levels (Hersbach et al., 2020 [21]). The hourly values have been linearly interpolated to the TROPOMI orbit timestamp and re-gridded to a $0.1° \times 0.1°$ resolution.

2.3 CAMS real-time fields

The Copernicus Atmospheric Monitoring Service (CAMS) global near-real-time service provides analyses and forecasts for reactive gases, greenhouse gases and aerosols on 25 pressure levels with a horizontal resolution of $0.4° \times 0.4°$ and a temporal resolution of 3 hours (Huijnen et al., 2016 [22]). For the calculation of NO_x emissions from TROPOMI observations, we use CAMS concentration fields of nitrogen oxides (NO and NO_2) and hydroxyl radical (OH). We also use the CAMS temperature field T. NO and NO_2 concentrations are used to account for chemical processes that take place in polluted air. Anthropogenic activities produce mainly NO, which is transformed into NO_2 by reaction with ozone O_3. NO_2 is then photolyzed during the day, reforming NO (Seinfeld, 1989 [23]). This photochemical equilibrium between NO and NO_2 can be highlighted with the NO_x:NO_2 concentration ratio, whose value depends on local conditions, allowing to perform a conversion from NO_2 production to NO_x emissions. The reason for the use of OH is different. OH is the main oxidant that controls the ability of the atmosphere to remove pollutants such as NO_2 (Logan et al., 1981 [24]). It is mainly produced during daylight hours by interaction between water and atomic oxygen produced by ozone dissociation (Levy, 1971 [25]). In air that is directly influenced by pollution, another source of OH is due to a reaction between NO and HO_2. This reaction, referred to as the NO_x recycling mechanism, illustrates the nonlinear dependence of the OH concentration on NO_2 (Valin et al., 2011 [26]; Lelieveld et al., 2016 [27]). Since the OH lifetime is typically of less than a second, its concentration in the troposphere is very low and difficult to measure. As a consequence, global analyses, which estimate OH concentrations from other variable species (Li et al., 2018 [28]; Wolfe et al., 2019 [29]), provide a representation for OH concentrations with high associated uncertainties. Therefore, the CAMS OH concentrations are used here to account for the NO_2 oxidation to form nitric acid (HNO_3), whose representation is explained in Section 3.1. Finally, the temperature field is used to control variations in the kinetic parameters (Burkholder et al., 2020 [30]). The hourly values are also linearly interpolated to the TROPOMI orbit timestamp and re-gridded to a $0.1° \times 0.1°$ resolution.

2.4 Background removal

Detecting traces of anthropogenic emissions in TROPOMI NO_2 images is not a straightforward process. The NO_2 signal from a sparsely populated area or a small industrial facility may be covered by numerical noise or by the signal generated by natural NO_x emissions. In the absence of anthropogenic sources, TROPOMI observes NO_2 concentrations which constitute a tropospheric background of $\sim 0.5 \times 10^{15}$ molecules.cm^{-2}. At the global scale, this background is the result of different sources. In the lower troposphere, natural NO_x emissions are mostly due to fires and soil emissions (Yienger et al., 1995 [31], Hoelzemann et al., 2004 [32]). In the upper troposphere however, sources include lightning, convective injection and downwelling from the stratosphere (Ehhalt et al., 1992 [33]), but the factors controlling the resulting concentrations are poorly understood. According to state-of-art estimates, anthropogenic NO_x accounts for most of the emissions at the global scale, whereas natural emissions from fires, soils and lightning are less significant at the global scale and do not exceed a share of 35% combined (Jaeglé et al., 2005 [34]; Müller and Stavrakou, 2005 [35]), although associated errors can be very high. In eastern China, the non-anthropogenic share of total NO_x emissions is variable but does not exceed 20% (Lin, 2012 [36]). Egypt being a desertic region and not very conductive to lightning, we expect the share of those non-anthropogenic emissions to be smaller. To estimate anthropogenic NO_x emissions, it is therefore necessary to remove this share.

With an atmospheric lifetime of about a few hours, the presence of NO_2 is relatively short. Consequently, the majority of NO_2 is not transported far downwind from its sources. Thus, near-surface NO_2 concentrations are generally high over industrial facilities and densely populated areas that need to be identified. Because Egypt's population is almost entirely located along the Nile River and its delta, the study of NO_x emissions in this country cannot therefore be reduced to the study of a small number of point sources, as it would be the case for several other parts of the

Middle East region, and must be carried out in the form of a mapping of the country. Further discussion is provided in Section 3.3. To identify urban areas in Egypt, we use the Socioeconomic Data and Applications Center (SEDAC) GRUMP (Global Rural-Urban Mapping Project) data archive, which comprises eight global datasets, including a population density grid provided at a resolution of 30 arc seconds, with population estimates normalised for the year 2000 (CIESIN, 2019 [37]). We combine this database with field data giving the location of industrial facilities from energy-intensive industries in the region (data have been retrieved from the Global Energy Observatory for oil and gas-fired power plants, from IndustryAbout for aluminium and iron smelters, from the work of Elvidge et al., 2016 [38] for flaring sites, and from the work of Steven J. Davis and Dan Tong for cement plants; links at the end of this article).
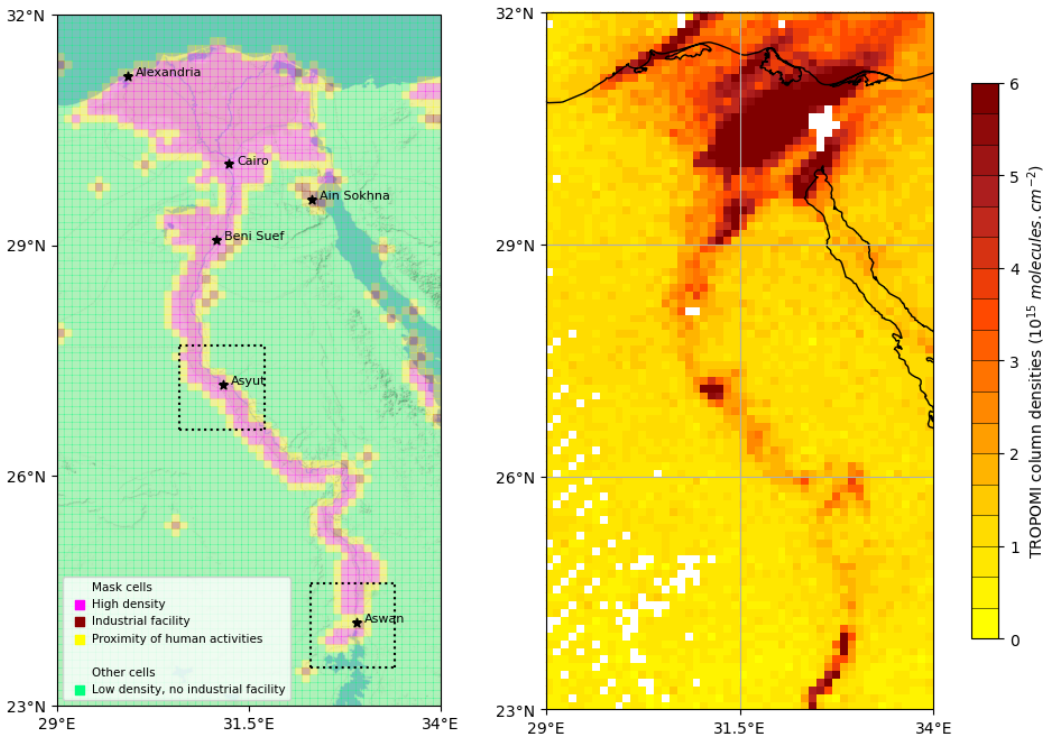

Figure 1: (left) Part of Egypt centered on Nile River. Within this domain, pink cells represent locations with an average human density above 100 hab.km^{-2}, brown cells represent locations with industrial facilities outside cities, and yellow cells represent locations in their vicinity. These cells constitute the mask used to calculate anthropogenic emissions. Outside this mask, green cells represent areas which do not correspond to any of the three criteria, considered to be void of human activity. Five larges cities in the country and the industrial area of Ain Sokhna are denoted with stars. Two smaller domains centered around the cities of Asyut and Aswan are represented with dotted lines; their use is presented in Section 4.6. (right) TROPOMI observation of NO$_2$ slant column densities above Nile valley on 3 January 2019. White pixels correspond to areas with low quality data ($q_a < 0.75$) or no data.

These datasets are used to remove the non-anthropogenic part of the NO$_x$ emissions signal. We conduct this removal by subtracting the mean emissions over areas without human activity from the mean emissions over industrial and densely populated areas. In order to perform this distinction between these two types of areas, our study is carried out using a mask within a $0.1° \times 0.1°$ grid. A grid cell is considered to be part of the mask if it has a population density higher than a threshold of 100 hab.km^{-2}, or if its centre is close to an industrial facility. Otherwise, the cell is considered to be part of the "background", i.e. outside the mask. In order to avoid any smearing that would correspond to abnormally high emissions outside urban and industrial centres (which can happen if the wind is poorly estimated), transition cells (in the immediate vicinity of the mentioned mask cells) are also considered to be mask cells. Figure 1 shows the distinction between mask cells and background cells on our domain in Egypt that lies between parallels 23°N and 32°N and meridians 29°E and 34°E. Most of the mask cells are located in the Nile area. Some mask cells are also found on the coast or in isolated parts in the desert. They correspond to remote industrial facilities, including major flaring sites, or sparsely populated industrial centres such as Ain Sokhna's industrial area. The domain comprises $n_m = 949$ mask cells and $n_b = 3692$ background cells. The mathematical description of the background removal is outlined in Section 3.4.

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2.5 Emission inventories

171 We compare TROPOMI-derived NO_x emissions to the Emissions Database for Global Atmospheric Research
172 (EDGARv5.0) for 2020 and the CAMS global anthropogenic emissions (CAMS-GLOB-ANT_v4.2) inventory released
173 in 2020. Both datasets provide $0.1° \times 0.1°$ gridded emissions for different sectors on a monthly basis. EDGARv5.0
174 emissions are based on activity data (population, energy production, fossil fuel extraction, industrial processes, agricul-
175 tural statistics, etc.) derived from the International Energy Agency (IEA) and the Food and Agriculture Organization
176 (FAO), corresponding emission factors, national and regional information on technology mix data and end-of-pipe
177 measurements. The inventory covers the years 1970-2015 and differs from the previous version EDGARv4.3.2 which
178 does not use splitting factors derived from the Energy Information Administration (EIA) data on fuel consumption of
179 coal, oil and natural gas for specific countries (Crippa et al., 2020 [39]). CAMS-GLOB-ANT_v4.2 is developed within
180 the framework of the Copernicus Atmospheric Monitoring Service (Granier et al., 2019 [40]). For this inventory, NO_x
181 emissions are based on various existing sectors in the EDGARv4.3.2 emissions from 2000-2012 which are used as a basis
182 for 2010 emissions and are extrapolated to the current year using 2011-2014 sector-based trends from the Community
183 Emissions Data System (CEDS) inventory (Hoesly et al., 2018 [41]). From one inventory to another, the names and
184 definitions of the sectors may vary. In EDGARv5.0 and CAMS-GLOB-ANT_v4.2, the emissions for a given country
185 are derived from the type of technologies used, the dependence of emission factors on fuel type, combustion conditions,
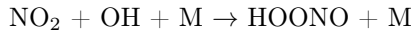186 as well as activity data and low resolution emission factors (Janssens-Maenhout et al., 2019 [42]).

187 **3 Method**

188 **3.1 Calculation of NO_2 production from TROPOMI observations**

189 As a first step, we use tropospheric NO_2 vertical column densities Ω_{NO_2} to derive top-down NO_2 production maps.
190 Vertical column densities Ω_{NO_2} are obtained from TROPOMI slant column densities using an air mass factor (AMF)
191 which is also provided by TROPOMI. Previous studies have shown that the use of the AMF is a source of structural
192 uncertainty in NO_2 retrievals (Boersma et al., 2004 [43]; Lorente et al., 2017 [44]). In polluted environments, this
193 source of uncertainty becomes non-negligible. Here, the AMF does not vary much temporally throughout the studied
194 period, and is around 1.6 for mask cells and around 1.9 for background cells. The difference between the two types of
195 cells is probably due to a different albedo between the urban environment and desert areas. Using the horizontal wind
196 $\mathbf{w}$, the NO_2 flux is given as $\Omega_{NO_2}\mathbf{w}$. The divergence of this flux can be added to the local time derivative $\frac{\partial \Omega_{NO_2}}{\partial t}$ to
197 balance NO_2 sources e_{NO_2} and sinks s_{NO_2} according to the continuity equation:

$$\frac{\partial \Omega_{NO_2}}{\partial t} + \mathrm{div}(\Omega_{NO_2}\mathbf{w}) = e_{NO_2} - s_{NO_2} \tag{1}$$

198 In steady state, the time derivative disappears and the mass balance is reduced to three terms. The NO_2 production
199 can thus be estimated by taking the combined effect of atmospheric transport losses and the different sinks. For the
200 transport term, we calculate numerical derivatives with a fourth-order central-finite difference scheme for each cell of
201 the domain. Moreover, since the local overpass time of TROPOMI occurs is in the middle of the day, NO_2 losses are
202 largely dominated by the three-body reaction involving NO_2 and OH (Seinfeld, 1989 [23]). Two channels have been
203 identified for this reaction (Burkholder et al., 2020 [30]), leading to the production of nitric acid HNO_3 and pernitrous
204 acid HOONO:
205 $NO_2 + OH + M \rightarrow HNO_3 + M$
206 $NO_2 + OH + M \rightarrow HOONO + M$
207 For the OH concentrations that are considered in this region (1-20 $\times 10^6$ molecules.cm^{-3}), the reactions above follow
208 first order kinetics. The total sink term can therefore be calculated as $s_{NO_2} = \Omega_{NO_2}/\tau$ with:

$$\tau = \frac{1}{k_{\mathrm{mean}}(T, [M]) \cdot [OH]} \tag{2}$$

209 τ appears here as the characteristic mixed lifetime of NO_2 in the atmosphere. The reaction rate k_{mean} characterises
210 the reactions between NO_2 and OH and depends on atmospheric conditions. Burkholder et al., 2020 [30] provide a
211 general expression of this rate as a function of both temperature T and total air concentration [M]. Note that although
212 this reaction rate accounts for both reactions with OH, the second channel is minor, because HOONO can be rapidly
213 decomposed back to NO_2 and OH in the lower troposphere. The value of k_{mean} therefore represents the total loss of
214 NO_2 due to OH, with a contribution of the HOONO forming reaction between 5 to 15% under atmospheric conditions

215 (Sander et al., 2011 [45]; Nault et al., 2016 [46]). Thus, the NO_2 production can be calculated as the sum of a transport
216 term and a sink term:

$$e_{NO_2} = \text{div}(\Omega_{NO_2}\mathbf{w}) + \Omega_{NO_2}/\tau \tag{3}$$

217 The treatment for NO_x removal is simplified here. NO_x concentrations are influenced by other sinks. Stavrakou et
218 al., 2013 [47] showed that the reaction between NO_2 and OH forming HNO_3 accounted for 26 to 64% of total NO_x
219 loss at the global scale. However, the features of the climate in Egypt during daytime hinder many other processes
220 to have a significant effect. The following NO_x sinks, which can be of notable importance at the global scale, are not
221 taken into account here:
222 - NO_2 deposition through the leaf stomata of vegetation. This sink can be significant in forested areas. In Egypt,
223 the leaf area index is very low, except in the croplands of the Nile Delta where it is comparable to that of southern
224 Europe or the western United States (Fang et al., 2019 [48]), for which the corresponding lifetime was of about 10-
225 100 h (Delaria et al., 2020 [49]), i.e. about an order of magnitude larger than the lifetimes calculated here. To our
226 knowledge, there are no studies focusing on the corresponding lifetimes for croplands, and we therefore do not take
227 this sink into account.
228 - NO_x oxidation by organic radicals to produce of alkyl and multifunctional nitrates (Sobanski et al., 2017
229 [50]). This sink increases with the concentration of volatile organic compounds (VOC), whose presence cannot be
230 excluded in Egypt. Different models have estimated low biogenic isoprene emissions in the region. These emissions
231 are concentrated at the level of the Nile and its delta (Guenther et al., 2006 [51]). They are certainly noticeable and
232 higher in summer than in winter, and contrast with the rest of the country, but they remain low compared to other
233 regions in the world. They are, for instance, about an order of magnitude lower than in the forested areas of the
234 eastern US, where the corresponding sink accounts for between 30% and 60% of the total NO_x sink (Romer Present
235 et al., 2020 [52]). Furthermore, at large NO_2 concentrations (compared to VOC concentrations), the share of this sink
236 in the total NO_x loss is weakened compared to that of HNO_3 (Romer Present et al., 2020 [52]). The effect of biogenic
237 emissions of VOC can therefore be considered minor. However, VOC emissions can also be of anthropogenic origin,
238 especially in urban areas, where they are difficult to estimate. To our knowledge, there is no study evaluating the
239 competition of the two sinks in Egypt or in a region with similar features and we therefore do not account for this
240 reaction in our calculations.
241 - NO reaction with HO_2 to produce nitric acid (Butkovskaya et al., 2005 [53]), whose yield is strongly enhanced in
242 presence of water vapour (Butkovskaya et al., 2009 [54]). Here, we neglect this reaction as the corresponding reaction
243 rate is lower by a factor 3 to 8 in dry conditions (Butkovskaya et al., 2005 [53]);
244 - NO conversion to NO_3, the latter being in thermal equilibrium with NO_2 and N_2O_5. This sink, which takes place
245 via heterogeneous processes, has a significant contribution during nighttime in the Mediterranean region (Friedrich et
246 al., 2021 [55]), is neglected at 13:30 when OH is close to its daily maximum;
247 - NO_2 reversible reaction with peroxyacetyl radical to produce peroxyacetylnitrate (Moxim et al., 1996 [56]). In
248 the Nile Delta region, PAN concentrations in the lower troposphere are significantly below the global average (Fischer
249 et al., 2014 [57]), suggesting a small yield. Moreover, its production peaks in the late afternoon and early evening
250 (Seinfeld, 1989 [23]). We therefore do not consider this sink in the representation of NO_x emissions at 13:30;
251 - NO_2 uptake onto black carbon particles (Longfellow et al., 1999 [58]). This uptake is of limited amount in the
252 Mediterranean region (Friedrich et al., 2021 [55]).

253 All these processes not being accounted for, the reaction between NO_2 and OH is the only sink that is considered
254 in our calculations to provide an indication of NO_x emissions. Section 4.7 details the consequences of not considering
255 these various minority sinks on the results.

3.2 Interpolation to daily average emissions

257 All parameters are evaluated at 13:30 local time, which means that the NO_2 production is calculated at the same
258 moment. In Egypt, the maximum and minimum electricity consumption are reached around 20:00 and 6:00 local time
259 respectively, and inter-daily consumption differences have been weakened by the increasing sales of air conditioning
260 and ventilation systems in the past decades (Attia et al., 2012 [59]). The daily load profiles provided by the National
261 Egyptian Electricity Holding Company show that the mean daily electricity consumption corresponds approximately
262 to the consumption at 13:30 in the country (EEHC, 2021 [60]). The difference between the two quantities being small
263 both in summer (about +2 to -3%) and winter (about -2 to -6%), we consider our inferred emissions as representative
264 of the average activity in Egypt at any time of the year. This assumes that electricity consumption dominates
265 the emissions of the country, or that the other emitting sectors have a similar daily profile. This can be justified.
266 According to CAMS-GLOB-ANT_v4.2, the power sector accounts for 50 to 60% of total NOx emissions in Egypt.

EDGARv5.0 presents a lower share (40 to 45% of total emissions). Moreover, for both inventories, the transport sector accounts for the majority of the remaining emissions. According to the traffic congestion index in Cairo (https://www.tomtom.com/en_gb/traffic-index/cairo-traffic/), the congestion level around 13:30 seems to be slightly higher than during the morning peak, but lower than the during night peak. Traffic emissions at this moment of the day could therefore be representative of the average traffic emissions as well.

3.3 Validation of CAMS OH concentration using line density calculations for Riyadh

When the transport term is integrated over large spatial scales, it cancels out due to the mass balance in the continuity equation between NO_2 sources and NO_2 sinks. At first order, the integration of the inferred emissions over the whole domain (of about 490,000 km^2) thus reflects chemical losses of the sink term. In this term, the NO_2 lifetime calculation involves the reaction rate k_{mean}, whose annual variability is low due to small changes in Egyptian midday temperatures throughout the year, and OH concentration, whose annual variability is highly marked. In Egypt, tropospheric OH concentrations, which are strongly correlated with solar ultraviolet radiation (Rohrer and Berresheim, 2006 [61]) and NO_x emissions, are higher in summer than in winter. To ensure an appropriate representation of the OH field by CAMS data, we select a large number of TROPOMI images characterised by a homogeneous wind field, in which we calculate the NO_2 lifetime according to Equation (2), where [OH] corresponds to the near-real-time CAMS data and k_{mean} is calculated with the formula from Burkholder et al., 2020 [30]. We compare this value with the lifetime determined by a method initially developed by Beirle et al., 2011 [62], and expanded by Valin et al., 2013 [63] by introducing a rotation of the image to refine the chemical lifetime. This method consists in fitting an exponentially modified Gaussian function (EMG) to NO_2 line density profiles. These profiles correspond to the integrated NO_2 columns along the mean wind direction in the pollution pattern and centered around the source. They are obtained by rotating TROPOMI images in the mean wind direction and using the values of the nearest columns in a 100 km^2 area. Line density profiles are generated on a span of 300 km. An example is given in Figure 3. Within the average profile, the NO_2 burden and lifetime can be derived from the parameters that describe the best statistical fit. The EMG model is expressed as follows (Lange et al., 2021 [17]):

$$\langle \Omega_{NO_2} \rangle (x|B, A, x_0, \mu, \sigma) = B + \frac{A}{2x_0} \exp\left(\frac{\mu - x}{x_0} + \frac{\sigma^2}{2x_0^2} \right) \mathrm{erfc}\left(-\frac{1}{\sqrt{2}} \left(\frac{x - \mu}{\sigma} - \frac{\sigma}{x_0} \right) \right) \tag{4}$$

Here, x is the distance in the downwind-upwind direction, B is the NO_2 background, A is the total number of NO_2 molecules observed in the vicinity of the point source, x_0 is the e-folding distance downwind, representing the exponential length scale of NO_2 decay, μ is the location of the apparent source relative to the centre of the point source, and σ is the standard deviation of the Gaussian function, representing the length scale of Gaussian smoothing. Using a non-linear least squares fit, we estimate the five unknown parameters: A, B, x_0, μ and σ. From the mean wind module w_{mean} in the domain, the mean effective NO_2 lifetime τ_{fit} can be estimated using the fitted parameters:

$$\tau_{fit} = \frac{x_0}{w_{mean}} \tag{5}$$

The geography of Egypt does not suit the method described here. The Egyptian population is contiguously concentrated along the Nile, which makes it difficult to define point sources isolated from human activity. Furthermore, large isolated cities such as Alexandria or Suez are too close to the coast for the wind to be considered homogeneous. We therefore use the city of Riyadh, Saudi Arabia (24.684°N, 46.720°E) to perform the comparison between the CAMS-induced lifetime and the fit-obtained lifetime. Riyadh has been the focus of anterior studies (Valin et al., 2013 [63], Beirle et al., 2019 [15]), and is particularly suitable for several reasons. Firstly, Riyadh is a city within the latitudinal extend of Egypt (1600 km from Cairo) with a climate which is similar to the typical Egyptian climate. Secondly, NO_2 tropospheric columns over Riyadh are high ($\sim 9 \times 10^{15}$ molecules.cm^{-2}), leading to retrievals with a high signal-to-noise ratio. Thirdly, Riyadh is far from the coast, and its flat terrain makes the surrounding wind fields rather homogeneous during most of the year.

As the fitting algorithm is very sensitive to any disturbance that might be induced by NO_2 production from other point sources, it is necessary to identify heavy industrial facilities in the area. Riyadh is also an industrial area, with several power plants located close to the city centre. Figure 2 shows the location of the most important emitters in the region, which include five gas-fired power plants (PP7, PP8, PP9, PP10 and PP14), one oil-fired power plant (PP4) and one cement plant (CP). The five gas power plants, with a total capacity of more than 10 GW, are located in the periphery of the city. These power plants are sufficiently far away from the city centre for TROPOMI to distinguish their own emissions from those of Riyadh's centre with a resolution of 0.1° × 0.1°, which is not the case for CP and PP4 which are located in the city centre. It is therefore appropriate to restrict the study of NO_2 patterns over Riyadh

to days for which the emissions from the city centre and those from the gas power plants do not overlap. This is the case when the wind blows steadily and homogeneously in a north-south direction. Within about 150 km around the city centre, we thus calculate the average wind given by ERA5 and consider the observation as reliable if the mean angle $\langle\theta\rangle$ of the observations deviates by less than 40° from the north or the south, with a standard deviation σ_θ of less than 36°. This condition generally leads to a selection of observations with large wind speeds, low winds speeds being often associated with more variable directions. This ensures the NO_2 decay to be dominated by chemical removal and not by the variability of the wind (Valin et al., 2013 [63]). Finally, we select observations with clear-sky conditions downstream of the flow (with at least 80% downstream cells with $q_a > 0.75$).
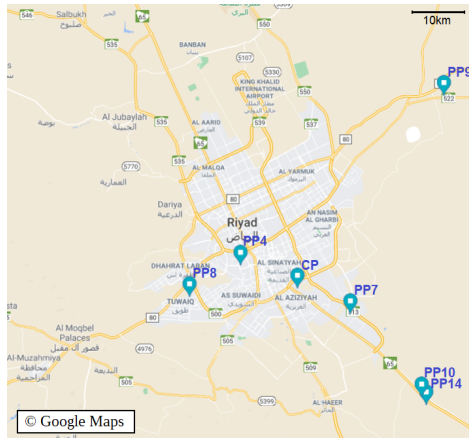
Figure 2: Map of Riyadh's city centre with the surrounding power plants (PP4, 7, 8, 9, 10 and 14) and cement plant (CP). The map has been extracted from Google Maps.

Our $0.1° \times 0.1°$ gridding ensures that retrieved lifetimes are governed by physical decay of NO_2 and not an artifact of the spatial resolution (Valin et al., 2011 [26]). The fitting procedure is very sensitive to the wind direction. Instead of manually correcting the ERA5 wind field for individual NO_2 patterns, the curve fitting is performed for every sample with three different rotation angles, corresponding to the wind direction with a correction of $-10°$, $0°$ or $10°$. A record is kept if one of these three fits leads to a correlation with the corresponding NO_2 line density whose coefficient is greater than 0.97. Among the remaining samples, we keep those with a value of τ_{fit} greater than 1.0 hr (considered sufficiently high to be relevant). An example of curve fitting is given in Figure 3.
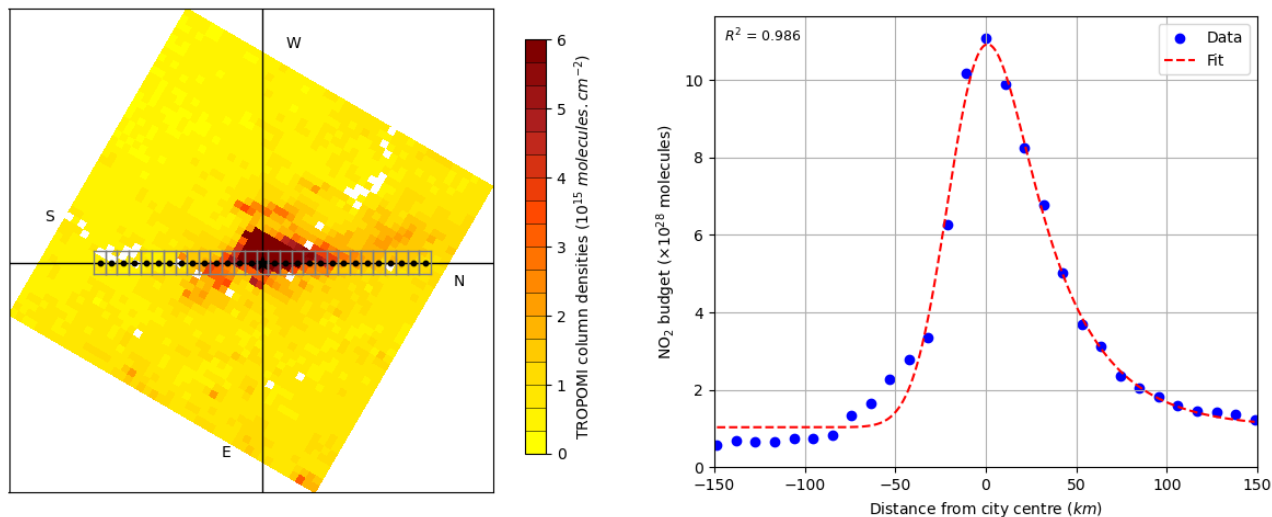

Figure 3: Estimation of the NO_2 lifetime from a pattern above Riyadh on 11 March 2020: (left) NO_2 plume rotated with its wind direction around the source (star) to an upwind-downwind pattern. Grey boxes centered around black points indicate the extent of the spatial integration of NO_2 columns to obtain the NO_2 line density. Values of cardinal points are noted in black. (right) Corresponding line densities (points) representing the downwind evolution of NO_2 as function of the distance to Riyadh's city centre, and the corresponding fit according to the exponentially modified Gaussian function (dashed line).

The phenomena under study here take place in the planetary boundary layer (PBL), which in this region has a midday height of about 2 km (Filioglou et al., 2020 [64]). TROPOMI observations only provide information on the total NO$_2$ content of the tropospheric column, without providing information on the vertical distribution of concentrations. Extracting emissions from concentrations therefore requires a selection on the height at which wind, temperature and OH data are taken. Lama et al., 2020 [19] and Lorente et al., 2019 [16] conducted similar studies using the boundary layer average wind, while Beirle et al., 2019 [15] chose a vertical level of about 450 m above ground. Because vertical transport of NO$_x$, which is emitted mainly from combustion engines and industrial stacks, is generally minor compared to horizontal transport, NO$_x$ is confined to the first few hundred metres above ground level. Using PBL-averaged data poses a problem of consistency as wind, temperature and OH concentration values significantly vary within the PBL. As a consequence, we compare the CAMS-induced lifetime τ and the fit-induced lifetime τ_{fit} using the parameters ($\mathbf{w}$, [OH] and T) at two different vertical levels: a medium level $\mathcal{A}$ at 925 hPa (about 770 m above ground level), and a bottom level $\mathcal{B}$ at 987.5 hPa (about 210 m). These levels are interpolated from four and two ECMWF or CAMS consecutive pressure levels respectively (1000-850 hPa for level $\mathcal{A}$ and 1000-975 hPa for level level $\mathcal{B}$). Most mask cells having an altitude between 0 and 150 m, the corresponding pressure variations are small (up to $\sim$ 16 hPa), which allows us to neglect the effects of topography on the position of pressure levels. Figure 4 sums up the selection method for the comparison of lifetimes.
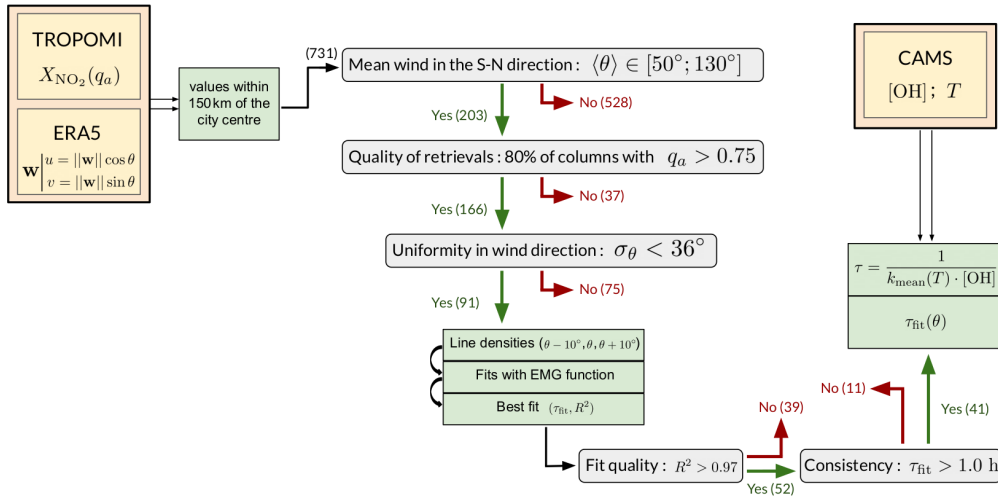

Figure 4: Selection method for NO$_2$ patterns over Riyadh. Datasets (yellow-orange) are used to calculate the quantities (light green) that are submitted to different tests (grey). 731 patterns are progressively conserved (green arrows) or rejected (red arrows). At each stage, the number of conserved or rejected patterns are noted within brackets (the value is only given for calculations performed at level $\mathcal{B}$). This selection process compares the lifetimes estimated by the EMG function fitting with TROPOMI line density profiles to the lifetimes calculated according to Equation (2) with CAMS data.

3.4 Calculation of anthropogenic NO$_x$ emissions and comparison with inventories

We calculate NO$_x$ emissions on the entire domain from NO$_2$ production by using CAMS NO and NO$_2$ concentrations. These are not intended to replace TROPOMI observations; they are used to apply the concentration ratio [NO$_x$]/[NO$_2$] = ([NO]+[NO$_2$])/[NO$_2$] to account for the conversion of NO$_2$ to NO and vice versa. As diurnal NO concentrations in urban areas generally range from 10 to 150 ppb (Khoder, 2008 [65]), the characteristic stabilization time of this ratio never exceeds a few minutes (Graedel et al., 1976 [66]; Seinfeld and Pandis, 2006 [67]). This time being lower than the order of magnitude of the inter-mesh transport time (about 30 min considering the resolution used and the mean wind module in the region), we can reasonably neglect the effect of the stabilization time of the conversion factor on the total composition of the emissions and treat each cell of the grid independently from its neighbours. Beirle et al., 2019 [15] found an annual average of 1.32 for this conversion factor, but CAMS data shows small deviations from this value over Egyptian urban areas. We therefore calculate NO$_x$ emissions for each cell of the domain as follows:

$$e_{\text{NO}_x} = \frac{[\text{NO}_x]}{[\text{NO}_2]} e_{\text{NO}_2} \tag{6}$$

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For convenience, quantities $\frac{[\mathrm{NO_x}]}{[\mathrm{NO_2}]}\mathrm{div}(\Omega_{\mathrm{NO_2}}\mathbf{w})$ and $\frac{[\mathrm{NO_x}]}{[\mathrm{NO_2}]}\Omega_{\mathrm{NO_2}}/\tau$ represent the respective contributions of the transport and the sink terms to total $\mathrm{NO_x}$ emissions. We finally obtain the emissions related to human activity $E_{\mathrm{NO_x}}$ by removing the arithmetic mean value of $\mathrm{NO_x}$ emissions above background cells from total emissions:

$$E_{\mathrm{NO_x}} = e_{\mathrm{NO_x}} - \frac{1}{n_b}\sum_{i=1}^{n_b} e_{\mathrm{NO_x},i} \tag{7}$$

These removed emissions are linked to the $\mathrm{NO_2}$ background estimated by TROPOMI, and do not correspond to anthropogenic emissions. They provide the value of what must be substracted from the estimates to obtain emissions related to human activity. Such a removal assumes that other processes involved in $\mathrm{NO_x}$ budgets lead to similar emissions inside and outside the mask, which is not evident, as the majority of background cells are located in the desert or the ocean while the majority of mask cells are located near the Nile River. However, as the processes involved in natural $\mathrm{NO_x}$ sources lead to emissions much smaller than anthropogenic emissions in polluted areas, we neglect this difference in the following calculations. An alternative would be to calculate an average concentration for the background cells and subtract the corresponding value from the column densities before calculating the emissions. This would pose further reliability problems: for instance, very high $\mathrm{NO_2}$ concentrations could appear outside the mask due to wind transport (an example is shown on Figure 1). They would lead to an overestimation of the $\mathrm{NO_2}$ background and thus to an underestimation of the anthropogenic emissions.

Neglecting the part of the country that lies outside the domain, total emissions from the anthropogenic activity of Egypt can then be obtained by integrating $E_{\mathrm{NO_x}}$ on the whole domain. For robust statistics, these derived emissions can be averaged monthly, enabling a month-by-month comparison with bottom-up inventories. The linearity of the averaging processes ensures the interchangeability of temporal and spatial averages. A monthly average is relevant because it aggregates enough data to limit the high inter-day variability due to changing wind patterns or differences between week days and week-ends. In addition, it enables the study of monthly $\mathrm{NO_x}$ emission profiles which reflect changes in human activities throughout the year due to temperature changes, economic constraints and cultural norms.

4   Results and discussion

4.1 Line densities and $\mathrm{NO_2}$ lifetime

We compare the results of the TROPOMI line densities fits for Riyadh to the lifetime calculated by Equation (2) using CAMS OH data. The two years of TROPOMI observations (from November 2018 to November 2020) provide a wide variety of profiles. For level $\mathcal{B}$, Figure 4 also provides the number of samples that are being kept at each stage of the process. Of the 731 observations available, 203 have a wind direction in the cone with a north-south orientation with an aperture of $40°$ (i.e. between $340°$ and $20°$ or between $160°$ and $200°$). Of the remaining observations, 166 occured with a sufficiently clear sky to be retained. The criterion of weak variability for the wind direction brings to 91 the number of observations that are kept by the method. On these 91 observations, the line density profiles are calculated and the fits applied. According to Equation (5), the lifetime is calculated using the mean wind module around the point source. The two lifetimes are calculated with the parameters taken at the medium level $\mathcal{A}$ or at the top level $\mathcal{B}$. Of the 91 fits obtained, 51 are of high quality (correlation coefficient between fit function and line density profile greater than 0.97) for level $\mathcal{A}$ and 52 for level $\mathcal{B}$. 39 of these fits lead to a lifetime τ_{fit} greater than 1.0 h for level $\mathcal{A}$ and 41 for level level $\mathcal{B}$. All remaining samples correspond to atmospheric conditions with moderate to fast winds, with a module ranging between 2 and 11 m/s (with an average of 5.9 m/s) for level $\mathcal{A}$ and between 3 and 8 m/s (with an average of 5.4 m/s) for level $\mathcal{B}$. These lifetimes are compared to the corresponding lifetimes obtained from CAMS data in Figure 5, which is divided into seasons for a more convenient comparison. The use of level $\mathcal{A}$ leads to notable underestimations of the $\mathrm{NO_2}$ lifetime in autumn compared to the lifetime calculated with the fitting method. This same level leads to lifetime overestimations in winter. This trend is not found with the use of level $\mathcal{B}$, which leads to a better reproduction of the lifetimes calculated with the fitting method for the available seasons. Figure 5 shows a linear regression between the two calculated lifetimes. The results are scattered, with a correlation coefficient higher for level $\mathcal{B}$ (0.408) than for level $\mathcal{A}$ (0.220). When the intercept of the regression line is forced to zero, the resulting slope is closer to 1 for level $\mathcal{B}$ (0.998) than for level $\mathcal{A}$ (1.071).
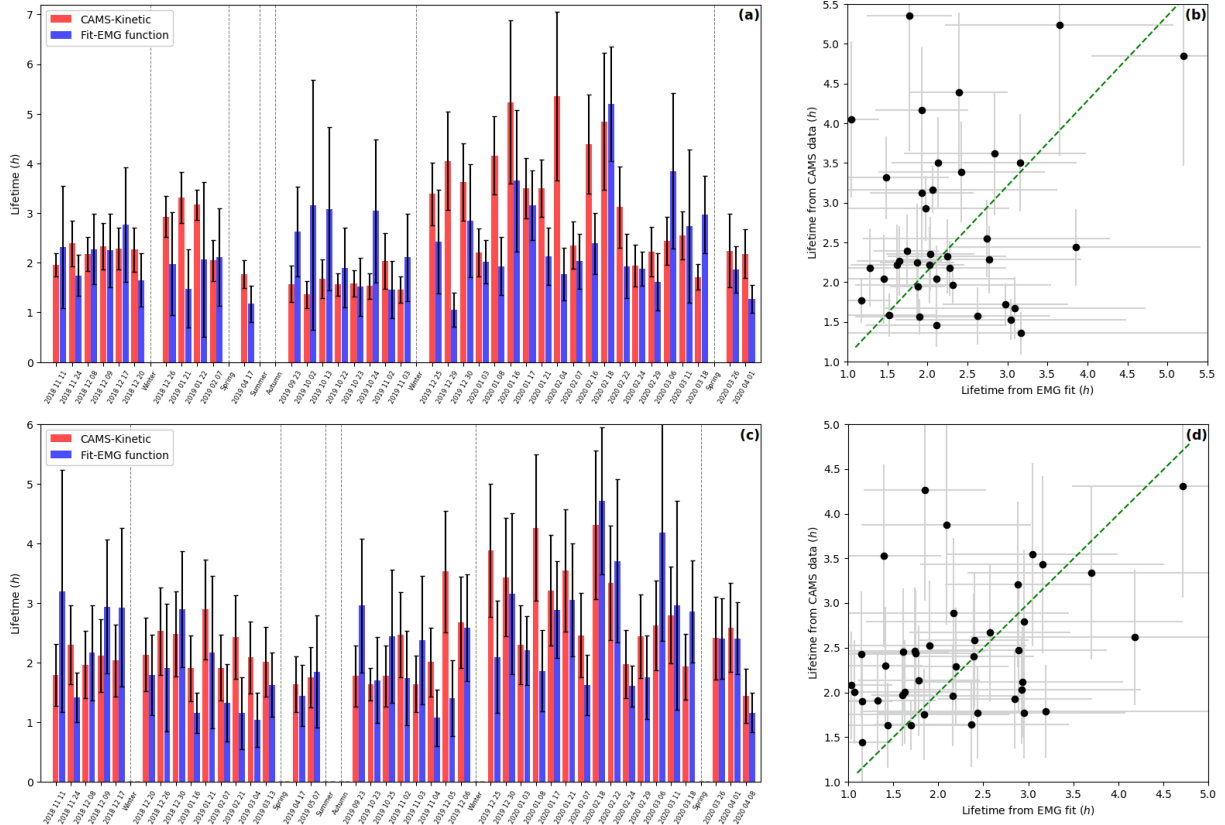
Figure 5: (left) Comparison between CAMS-derived NO_2 lifetimes and lifetimes from NO_2 line density fittings with EMG function above Riyadh city centre, for level $\mathcal{A}$ (a) and $\mathcal{B}$ (c). The samples presented correspond to patterns in clear-sky conditions for which the mean wind is in the north-south direction with a low variance, and for which the correlation between line density profile and fit gives a correlation coefficient of more than 0.97 and a lifetime of more than 1.0 h. NO_2 patterns do not have these conditions during the summers of 2019 and 2020. Dashed lines separate the groups of observations by season. (right) Comparison between the two calculated lifetimes for level $\mathcal{A}$ (b) and $\mathcal{B}$ (d). A linear regression with an intercept forced to be zero is displayed with a green dashed line.

Level $\mathcal{B}$ is therefore the one that leads to the best match between the lifetime calculated with Equation (2) and the lifetime calculated from line densities. The results that are presented in the following sections (except for Section 4.3) are therefore results of calculations performed with parameters ($\mathbf{w}$, [OH], T and $[NO_x]/[NO_2]$) estimated at level $\mathcal{B}$. Nevertheless, it should be noted that no summer observations were included in the comparison. The main reason for this is the wind direction: of the 188 summer days observed, 178 have a mean wind direction outside the north-south cone over central Riyadh. On the remaining ten days (one for summer 2019 and nine for summer 2020), the ERA5 wind direction is too variable for the fit to be considered relevant, or the fit results in a correlation coefficient below 0.97. It is not clear how correctly the NO_2 lifetime would be calculated during both summer periods by Equation (2). With OH concentrations being the main driver of this lifetime, we cannot assess the relevance of the representation of OH concentrations by CAMS data during summer days in the study.

4.2 Mapping of Egypt's NO_x emissions

The top-down emission model is then applied to the Egyptian domain with CAMS OH concentration and temperature fields for lifetime calculations. For each cell, NO_x emissions are calculated according to Equation (6), resulting in a mapping of Egypt's emissions. The obtained values are averaged monthly from November 2018 to November 2020. Figure 6 shows a composition of the emissions map with the transport term and the sink term for the months of January and July 2019. The corresponding anthropogenic emissions, calculated according to Equation (7), are added. The Nile appears on transport term maps: the divergence calculation complies with what is expected from a line of emitters, i.e. a clear separation of zones of positive divergence from zones of negative divergence with a separation line corresponding to the course of the river. The fact that areas of negative and positive divergence are respectively located to the east and the west of the river is the result of the zonal component of the wind being positive most of the time. Some point sources like Cairo, Alexandria, Asyut or Aswan are easily identifiable. The sink term, directly proportional to the TROPOMI column densities, also highlights these cities. However, unlike the transport term,

which has a similar spatial pattern from month to month, the sink term is clearly stronger in summer than in winter. This is mainly due to a higher lifetime in winter than in summer (4.94 h on average in January 2019 and 2.62 h in July 2019) while the average TROPOMI NO_2 concentrations are slightly higher during winter (1.071 $\times 10^{15}$ molecules.cm^{-2} for January 2019 and 0.899 $\times 10^{15}$ molecules.cm^{-2} for July 2019). Over the whole domain, the mean transport term varies throughout the studied period between -0.014×10^{15} molecules.cm^{-2}.h^{-1} (December 2019) and 0.015×10^{15} molecules.cm^{-2}.h^{-1} (May 2020). Thus, it hardly contributes to the NO_x emission budget, the mean chemical sink term alone varying between 0.223×10^{15} molecules.cm^{-2}.h^{-1} (December 2019) and 0.534×10^{15} molecules.cm^{-2}.h^{-1} (September 2020).
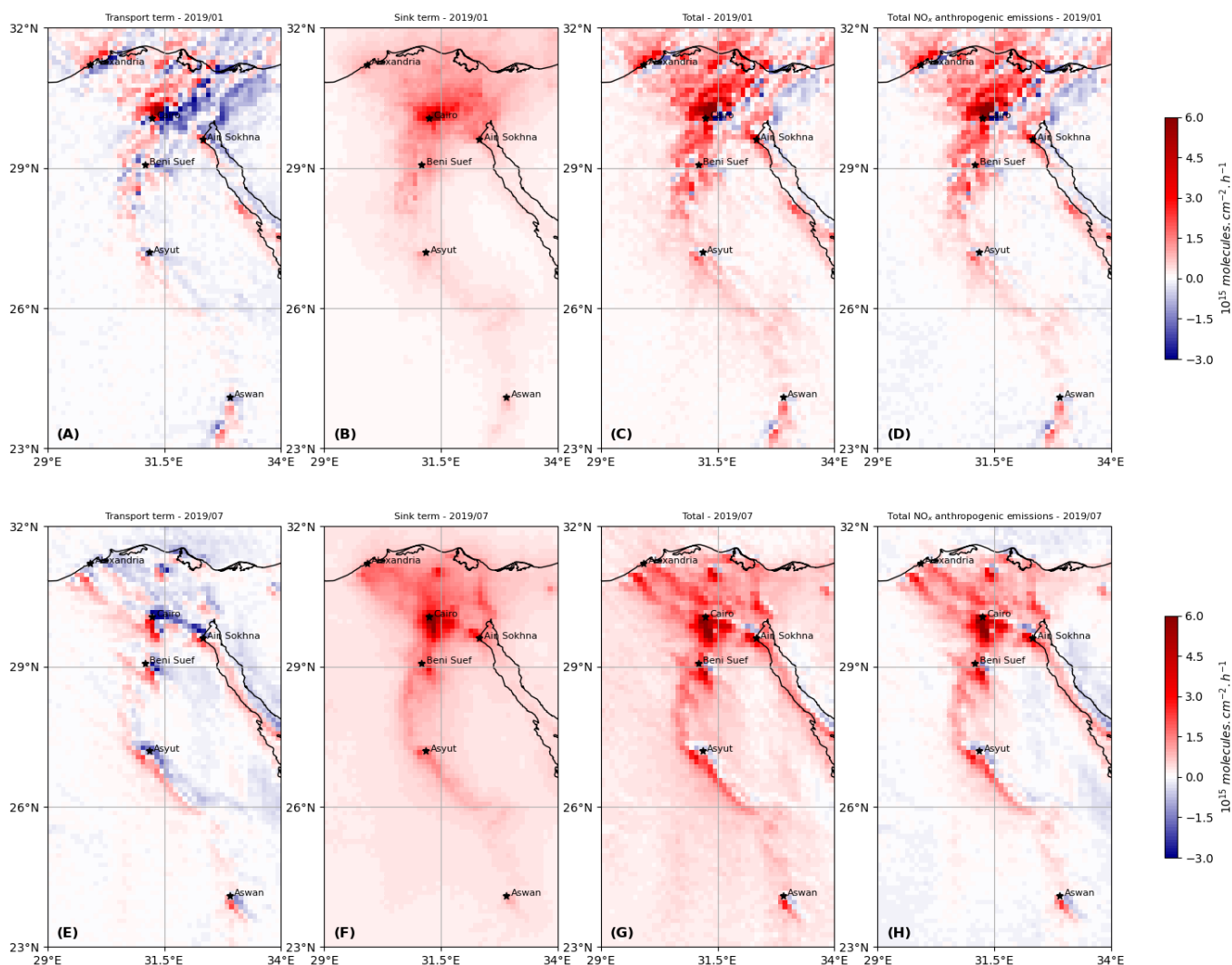
Figure 6: NO_x emissions above most of Egypt's territory: (top) transport term (A), sink term (B), resulting surface emissions (C), and the corresponding anthropogenic emissions after non-anthropogenic background removal (D) for January 2019. (bottom) transport term (E), sink term (F), resulting surface emissions (G), and the corresponding anthropogenic emissions after non-anthropogenic background removal (H) for July 2019.

Several cities in the country appear as the main emitters of the country, such as Cairo, Alexandria, Beni Suef, Asyut or Aswan. The industrial area of Ain-Sokhna, located southwest of Suez, also appears as a main emitter. Table 1 compares the monthly values for the sink term and the absolute value of the transport term above five major cities of the country, with populations ranging from 200,000 to 20 million inhabitants, as well as Ain-Sokhna's area. The mean values for TROPOMI column densities are also provided. According to the results, the capital city of Cairo is by far the largest emitter in the country, largely due to its large population, resulting in high traffic emissions, but also to its intensive industrial activity. Alexandria, the country's second largest city, is not necessarily the second largest emitter, as its emissions are comparable to those of smaller cities such as Beni Suef or Asyut. However, the three cities concentrate a large amount of industrial activity: Alexandria hosts several oil and gas power plants and a small

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number of cement factories, while Beni Suef is close to several oil and gas power plants and hosts several flaring sites. Similarly, the city centre of Asyut is close to three oil and gas-fired power plants and a cement factory. This seems to indicate that industrial activity might be the main cause of NO_x emissions differences between these cities, before population size. This explains why NO_x emissions from these three cities are comparable to those of the industrial area of Ain Sokhna, which includes several cement facilities, iron smelters and oil and gas plants. It might also explain why Aswan, which has a population that is comparable to Beni Suef or Asyut, but which does not have any major industrial site, has slightly lower emissions than the two other cities. An additional analysis of the differences between Asyut and Aswan is provided in Section 4.6. Finally, the Gulf of Suez displays relatively large emissions, which might be attributed to the shipping sector, the region being a major gateway for international trade. Because it also hosts several flaring sites, these emissions might also be due to the oil and gas extraction activity.

City	Population density	Jan. 2019			Jul. 2019		
		Ω_{NO_2}	Transport	Sink	Ω_{NO_2}	Transport	Sink
	(habitants per square kilometer)	$(\mathcal{M}_{NO_2}.cm^{-2})$	$(\mathcal{M}_{NO_x}.cm^{-2}.h^{-1})$		$(\mathcal{M}_{NO_2}.cm^{-2})$	$(\mathcal{M}_{NO_x}.cm^{-2}.h^{-1})$	
Cairo	18,064	9.415	2.903	3.684	5.618	2.022	4.879
Alexandria	9,133	3.034	1.179	0.975	1.674	0.410	1.421
Asyut	1,644	1.708	0.679	0.718	2.137	1.236	1.520
Aswan	319	0.976	0.182	0.473	0.871	0.308	0.523
Beni Suef	2,056	2.950	0.548	1.080	2.321	0.428	1.591
Ain Sokhna	5	3.133	1.256	1.115	2.561	1.346	1.757

Table 1: Comparison between the transport term and the sink term above different cities in Egypt, as well as the industrial region of Ain Sokhna located 45 km southwest of Suez for January and July 2019. TROPOMI vertical NO_2 columns, NO_x emissions and population densities correspond to average values within 18 km from the city centre. Unit $\mathcal{M}$ stands for a quantity of 10^{15} molecules (NO_2 or NO_x).

Although these cities and areas can be described as high-emission sites, the term responsible for these emissions differ from one site to the other. Figure 7 shows the contribution of the transport term (taken in absolute value) to total emissions for January and July 2019. Because wind fields are relatively homogeneous along the Nile on spatial scales of less than 100 km, NO_2 concentration gradients perceived by TROPOMI in the region mainly contribute to the increase of the transport term which can reach similar values as the sink term. However, it is never significantly higher than the sink term: due to a spread of the emissions over large urban areas, the behaviour of these cities is therefore different from that of a point source for which the transport term would be very high (Beirle et al., 2021 [68]).
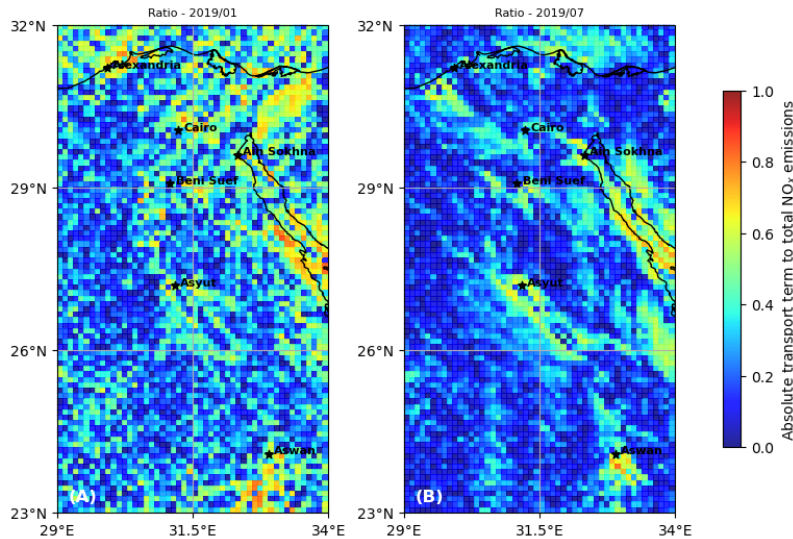

Figure 7: Share of the absolute value of the transport term in the sum of the sink term and the absolute value of the transport term above most of Egypt's territory, indicating the local importance of the transport term in NO_x emissions above mask cells. The average of this ratio is shown for January 2019 (A) and July 2019 (B).

Desert areas such as the Libyan Desert, the Eastern Desert and the Sinai region, (located respectively to the west, east and northeast of the Nile) show a very low value for the transport term compared to the sink term, due to the

homogeneity of both the wind field and the detected NO$_2$ concentrations in these areas. In the case of the Gulf of Suez, the transport term can be 1 to 2 times higher than the sink term, which varies between 0.4 and 1.2 $\times 10^{15}$ molecules.cm^{-2}.h^{-1}. Those values are slightly higher than the average emissions above background cells areas due to the sink term (about $0.2 - 0.6 \times 10^{15}$ molecules.cm^{-2}.h^{-1}), but remain quite low compared to the emissions in large cities. This relative predominance of the transport term is explained by a visible gradient of the TROPOMI NO$_2$ column densities. The region thus acts as a very thin line of emitters. Nevertheless, this predominance might also be partly due to a poor representation of the wind field. The low resolution of ERA-5 (about 26 km in this region, which is the same order of magnitude as the width of the channel) may misrepresent the wind near the coast, creating artificial gradients.

4.3 Vertical analysis

Here we investigate the influence of the choice of the vertical level in the representation of the model parameters. This influence is of considerable importance, as NO$_x$ sources in urban areas can be located at different altitudes. For instance, emissions from the road sector from tailpipes are located at ground level, whereas NO$_x$ from power plants and industrial facilities can be emitted from stacks, which are usually located between 50 and 300 m above the ground. Section 4.1 results showed that level $\mathcal{B}$ was more appropriate for the representation of the NO$_2$ lifetime. This level is therefore chosen as a reference for the comparison. We study the effect of a transition from level $\mathcal{B}$ to level $\mathcal{A}$ for each of the 3 parameters involved in the representation of the sink term, i.e. temperature T, hydroxyl radical concentration [OH] and concentration ratio [NO$_x$]/[NO$_2$]. The results for the averages over mask cells and background cells are given for the months of January, April, July and October 2019 in Table 2. As the wind field is only involved in the transport term whose spatial integration nearly leads to zero, the influence of this parameter is not studied.

		Sink term (10^{15} molecules.cm^{-2}.h^{-1})							
level $\mathcal{B}$ (987.5 hPa)	level $\mathcal{A}$ (925 hPa)	Jan. 19 (MASK)	Jan. 19 (BKGD)	Apr. 19 (MASK)	Apr. 19 (BKGD)	Jul. 19 (MASK)	Jul. 19 (BKGD)	Oct. 19 (MASK)	Oct. 19 (BKGD)
T, [OH], $\frac{[NO_x]}{[NO_2]}$	-	0.859	0.253	1.072	0.345	1.125	0.376	0.932	0.277
[OH], $\frac{[NO_x]}{[NO_2]}$	T	0.899 (+4.7%)	0.264 (+4.2%)	1.127 (+5.2%)	0.361 (+4.6%)	1.185 (+5.3%)	0.394 (+4.9%)	0.887 (+4.8%)	0.264 (+4.5%)
T, $\frac{[NO_x]}{[NO_2]}$	[OH]	0.769 (-10.5%)	0.219 (-13.6%)	1.013 (-5.5%)	0.324 (-6.0%)	1.129 (+0.4%)	0.375 (-0.3%)	0.853 (-8.5%)	0.251 (-9.5%)
T, [OH]	$\frac{[NO_x]}{[NO_2]}$	0.872 (+1.6%)	0.257 (+1.4%)	1.094 (+2.1%)	0.352 (+2.0%)	1.143 (+1.6%)	0.383 (+1.9%)	0.904 (+3.1%)	0.271 (+2.2%)

Table 2: Analysis of the effect of a vertical change in the parameters used to estimate the sink term in NO$_x$ emissions: temperature, hydroxyl radical concentration, and NO$_x$:NO$_2$ concentration ratio. The comparison is conducted between the estimated quantities for level $\mathcal{B}$ and level $\mathcal{A}$. The comparison is conducted for mask cells (MASK) and background cells (BKGD) for four months of the year 2019. Values within brackets represent the variation from the base case for which all quantities are estimated at level $\mathcal{B}$.

The transition to level $\mathcal{A}$ generally results in a decrease in temperature, leading to an increase in the reaction rate k_{mean} and thus an increase in the emissions from the sink term. This transition has only a small influence on the total NO$_x$ emission estimates, with mask and background cells emissions increasing by 4 to 6%. The influence of OH goes in the opposite direction: its concentration decreases with altitude, weakening the sink term. This weakening is particularly visible during winter months, for which the emissions are lower by up to 14%. In summer however, the effect is hardly noticeable. Finally, the influence of the NO$_x$:NO$_2$ ratio is negligible on the NO$_x$ emission estimates. Thus, the transition to level $\mathcal{A}$ results in an increase in the sink term of 1 to 4%, due to a decrease in both concentrations of NO and NO$_2$ with respect to the vertical but with a greater decrease for NO$_2$. This vertical study confirms the crucial importance of the OH concentration for the accurate representation of NO$_x$ emissions. It appears here as an important driver of the sink term, which is much more sensitive to vertical differences than temperature or the NO$_x$:NO$_2$ concentration ratio.

4.4 Weekly cycle

In Egypt, the official rest day is Friday, and the economic activity of the country is *a priori* lower during this day than during the other days of the week. We therefore try to characterise this feature, by evaluating the weekly cycle of NO$_x$ emissions. We use the TROPOMI-inferred emissions to obtain averages per day of the week. We use the quality assurance q_a of TROPOMI retrievals to ignore the days for which more than 20% of the domain has low-quality data (this happens 43 times in 2018/2019 and 28 times in 2019/2020). Such a filtering avoids accounting for the days when

a large part of the urban and industrial areas are covered by clouds. However, it misses situations where small clouds are localised over large emitters, in which case the corresponding emissions are under-estimated. Figure 8 shows the resulting daily emissions for the period November 2018 - November 2019 and November 2019 - November 2020. NO_x emissions are expressed in mass terms as NO. A Friday minimum is observed, defining a weekly cycle. This trend is also observed for mean NO_2 column densities, for which no intra-weekly variation is observed. Over the 2018-2019 period, Fridays have average emissions of 0.811 ± 0.408 kt, which is lower than average emissions for the rest of the week, which reach 0.997 ± 0.533 kt. A similar trend is observed in 2019-2020, for which the average for Fridays is 0.704 ± 0.357 kt and the average for other days is 0.921 ± 0.449 kt. The difference in emissions between the two periods is due to smaller emissions in December 2019, January 2020 and February 2020 that are discussed in Section 4.5. On average, Friday emissions correspond to a ratio of 0.83:7 (i.e. a value of 0.83 after normalisation on the seven days of the week) for the entire domain. This result is consistent with the values obtained by Stavrakou et al., 2020 [69], who used TROPOMI data and another emission model to calculate a ratio of 0.71:7 for Cairo and 0.89:7 for Alexandria in 2017.
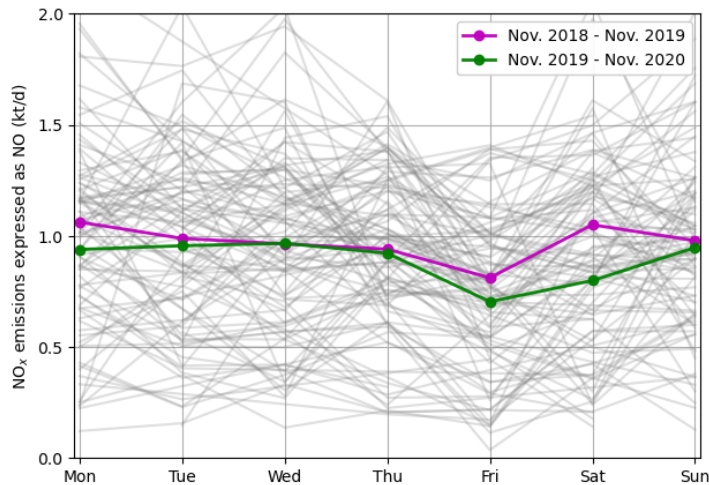

Figure 8: Weekly profiles of anthropogenic NO_x emissions for Egypt using TROPOMI observations in 2018-2019 (purple line) and 2019-2020 (green line). Thin grey lines represent individual weeks. Days for which less than 80% of the domain counts low quality observations ($q_a < 0.75$) are not represented.

4.5 Impacts of lockdown during COVID-19

The ongoing global outbreak of COVID-19 forced many countries around the world to implement unprecedented public health responses, including travel restrictions, quarantines, curfews and lockdowns. Such measures have helped to counter the spread of the virus and have, meanwhile, caused high reductions in global demand for fossil fuels (IEA, 2020 [70]). They also led to a fall in the levels of NO_2 and other air pollutants across the globe (Venter et al., 2020 [71]; Bauwens et al., 2020 [72]; Gkatzelis et al., 2021 [73]). To prevent the spread of COVID-19, Egyptian authorities ordered a partial lockdown from March 15th till June 30th 2020, closing all public areas (e.g. sport centres, nightclubs, restaurants and cafes) and suspending religious activities in all mosques and churches throughout the country. They also implemented more drastic measures such as a full lockdown during Easter (April 20th) and Eid (May 23rd to May 25th), before lifting some restrictions on June 1st (Hale et al., 2021 [74]). In addition to the effect of containment on the activity of the country, the global decline in consumption led to a drop in the production of certain industrial products.

Several studies have estimated the impact of these events on the air pollution levels in the urban centres of the country : from in-situ measurements, El-Sheekh et al., 2021 [75] estimated that NO_2 concentrations had dropped by 25.9% in Alexandria's city centre after the start of the lockdown on March 13th, while El-Magd et al., 2020 [76] used OMI retrievals to estimate a 45.5% reduction of NO_2 concentrations for the entire country during the spring compared to 2018 and 2019 average values. However, due to a changing lifetime of NO_2, reductions in the concentrations of NO_2 might not be entirely due to a decrease in NO_x emissions, which leads us to focus on the variation of NO_x emissions during this singular period. Using our top-down emission model, reductions in total NO_x emissions of 20.1%, 11.8% and 13.5% are observed for the months of March, April and May 2020 compared to the equivalent months in 2019. This drop of emissions in 2020 compared to 2019 calculated by the model also correspond to a decrease in observed

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NO$_2$ columns. No significant changes in OH concentrations seem to appear: on average, from 2019 to 2020, CAMS near-real-time data shows a decrease of 5.5% for OH concentration over the mask cells for the period March/April/May, while TROPOMI retrievals above mask cells show a decrease in NO$_2$ column densities of 21.6% over the same period. However, these effects observed for the months of March, April and May 2020 are not repeated in June 2020, for which emissions show an increase of 15.8% compared to June 2019. This rise is largely the result of an increase in the difference between average estimates inside and outside the mask. Indeed, emissions within the mask in June 2020 are higher than those of June 2019, due to an increase in TROPOMI concentrations above mask cells (+7.7%) while the NO$_2$ lifetime is almost unchanged (+3.3%). Emissions outside the mask varies in the opposite direction: a decrease in TROPOMI background concentrations (-5.4%) is observed while NO$_2$ lifetime increases strongly (+16.0%). This increase in June emissions seems to indicate that the lift on restrictions allowed a catch-up of the economic activity which was sufficiently strong to generate higher emissions in 2020 than in 2019.

4.6 Annual cycle and comparison to inventories

Here, we attempt to compare our TROPOMI-derived NO$_x$ emissions to emissions from CAMS-GLOB-ANT_v4.2 and EDGARv5.0 inventories. Figure 9 shows the total anthropogenic NO$_x$ emissions over the mask cells from November 2018 to November 2020, calculated according to Equation (7). As indicated in Section 3.2, the emissions, calculated at 13:30 local time, are representative of the average daily consumption in Egypt. The total calculated for each month therefore corresponds to the NO$_x$ production by human activities in the country. After aggregating the different sectors of activity, CAMS and EDGAR inventories directly provide the anthropogenic NO$_x$ emissions over the same domain. All NO$_x$ emissions are expressed in mass terms as NO. We note that the EDGAR inventory does not cover the period 2018-2020 (the last available year of the iventory is 2015). On Figure 9, EDGAR emissions corresponding to the period between November 2013 and November 2015 are displayed, i.e. the preceding 5 years compared to TROPOMI-derived emissions and CAMS estimates.
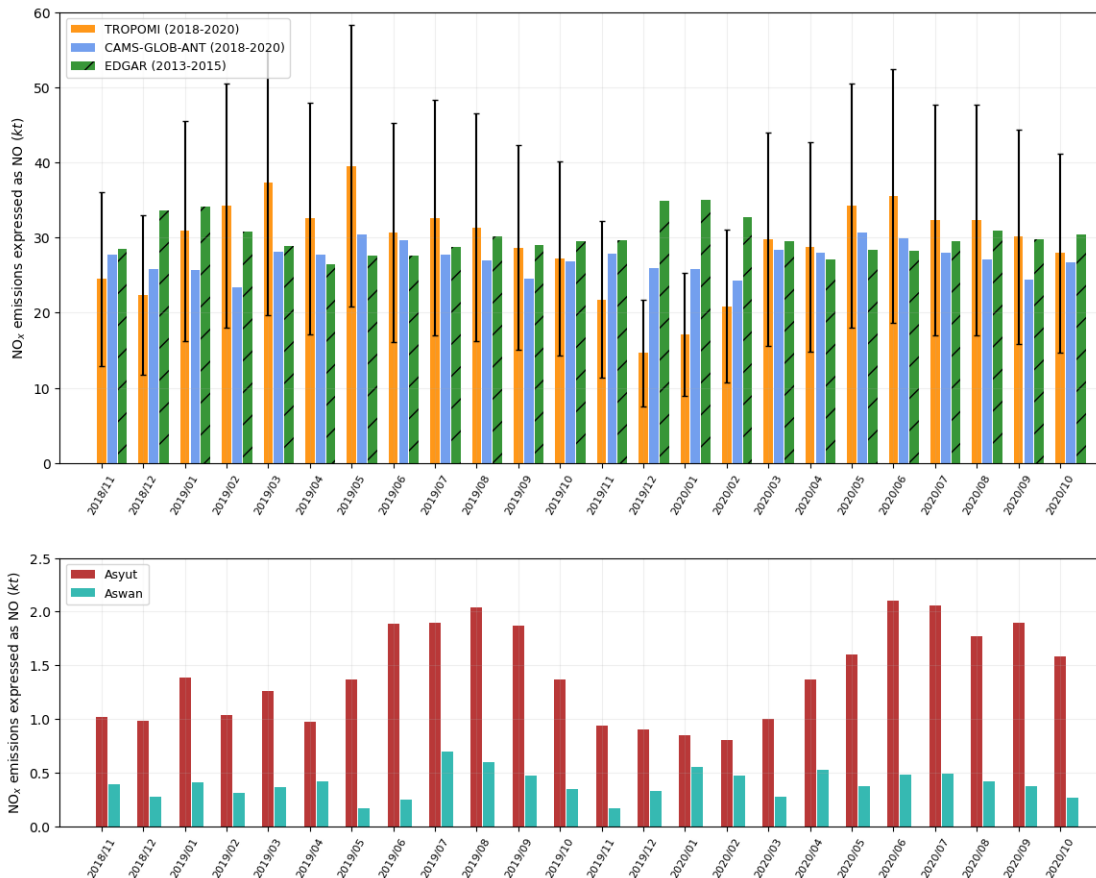

Figure 9: (top) Comparison of TROPOMI-derived anthropogenic NO$_x$ emissions in Egypt (light blue), with the corresponding emissions from EDGAR (green with stripes) and CAMS (yellow) inventories. EDGAR data is shown for comparison purposes and covers the years 2013-2015. (bottom) TROPOMI-derived anthropogenic NO$_x$ emissions for the cities of Asyut (dark red) and Aswan (light blue). The corresponding domains are displayed on Figure 1.

TROPOMI-derived emissions are higher than the CAMS inventory estimates. The top-down model estimates total emissions of 697.6 kt over the 24 months, which is 45.9 kt higher than CAMS for the same period (651.6 kt). This difference is primarily localized in the first 12 months, for which TROPOMI-inferred emissions are always higher than the inventories and show higher values in summer than during the rest of the year. The next 12 months show similar emissions in summer but much lower values in winter. In particular, the difference is significant in December 2019 and January 2020 (respectively 56.5% and 66.5% of CAMS levels). These emissions also contrast with other winter emissions, with a total of 31.7 kt for 2019-12/2020-01 against 53.3 kt for 2018-12/2019-01 and 57.7 kt for 2020-12/2021-01. In the computations, this drop for winter 2019/2020 is mainly due to a relatively low value of the OH concentration which reaches 4.61×10^6 molecules.cm^{-3} on average for December 2019 and January 2020, with 4.29×10^6 molecules.cm^{-3} above mask cells and 4.69×10^6 molecules.cm^{-3} over background cells. They were respectively 5.29, 5.74 and 5.18×10^6 molecules.cm^{-3} for the previous year (2018-12/2019-01) and 5.11, 4.90 and 5.16×10^6 molecules.cm^{-3} for the subsequent year (2020-12/2021-01). A decrease in tropospheric columns (-18.5% for mask cells and -7.6% for background cells compared to winter 2018/2019) also contributes to this drop. The accuracy of the inferred emissions for winter 2019/2020 can therefore be questioned.

At first sight, the annual variability of TROPOMI-inferred emissions, which describes a one-year cycle with higher emissions in summer, seems to be correlated with power emissions which dominate the use of fossil fuels in Egypt (Abdallah et al., 2020 [77]). These power emissions are due to the country's residential electricity consumption (Attia et al., 2012 [59]; Elharidi et al., 2013 [78]; Nassief, 2014 [79]). They also meet the needs of industry. Summer peaks in electricity consumption are mostly driven by temperature, as illustrated by the increasing sales of air conditioning and ventilation systems for several decades (Wahba et al., 2018 [80]). The use of air conditioning in cars, which requires an additional amount of fuel, could also contribute to the increase of NO$_\text{x}$ emissions in summer. To support this hypothesis, we use our model on two smaller domains centered around the two cities of Asyut and Aswan. The corresponding domains are displayed on Figure 1. Both cities have similar demographic features, with populations of about 467,000 and 315,000 inhabitants in 2021 and human densities of about 3,000 and 1,600 inhabitants per square kilometer respectively. However, their industrial features largely differ. There is no large fossil fuel-fired power plant in Aswan, where most of the electricity is produced by a hydroelectric dam, whereas Asyut counts three oil and gas power plants of various capacities (90, 650 and 1500 MW) in its urban area. Both cities have a cement plant, but the one in Asyut has a larger production capacity (5.7 Mt/yr in Asyut, 0.8 Mt/yr in Aswan). Our model is used following the same procedure as for the main domain. The background removal is done at the scale of the country. A seasonal cycle appears for Asyut, with a minimum for winter months and a maximum for summer months. This cycle seems slightly shifted from the one observed for the entire country, for which May emissions are as important as those of summer months. We also note that the decrease in emissions for winter 2019/2020 is less marked than for the emissions of the whole country, and of a similar value to the previous winter. This suggests that national NO$_\text{x}$ emissions are indeed lower during winter, but that the values obtained for winter 2019/2020 are particularly low. We also find that the seasonality of the emissions is more pronounced for the Asyut domain than for the country as a whole. The case of Aswan is different. Emissions within the corresponding domain are significantly lower than for Asyut. The signal-to-noise ratio being higher, it is difficult to characterise an annual cycle, but the results do not seem to indicate low emissions in winter and high emissions in summer. This identification of a seasonal cycle identical to that of the entire country for a city with several power plants, and the absence of such a cycle in a city without any, strengthens the hypothesis that the power sector plays a major role in Egyptian NO$_\text{x}$ emissions.

We note that some features of the industrial activities in the country might be counteracting this trend. For some sectors such as cement or steel, production is lower in summer, due to the physical wear experienced by workers due to heat, but also due to certain periods of leave. Given the importance of industrial activities in the production of NO$_\text{x}$ shown in Section 4.2, this aspect cannot be neglected. The transport sector could also counteract the observed trend: although the use of air conditioning in cars increase NO$_\text{x}$ emissions of the sector, the observed mean traffic in the country is higher between November and February and lower between June and August, especially in Cairo which gathers most of the population. In the absence of additional data, it therefore seems difficult to conclude on the amplitude of the seasonal cycle produced by our top-down model. This caution is all the more necessary as CAMS and EDGAR show seasonal cycles for NO$_\text{x}$ emissions, with different dynamics than those displayed by TROPOMI emissions: while the EDGAR inventory predicts a maximum of emissions in December or January and a minimum in April, the CAMS inventory shows two local maxima each year in May and November and two local minima in February and September. The amplitude of the corresponding cycles is much lower in those inventories, representing 14.2% of the average value for emissions estimates for EDGAR and 12.4% for CAMS. Those values must be compared to the amplitude displayed by TROPOMI-inferred emissions, for which the maximum/minimum ratio is about 1.8 if winter 2019/2020 is excluded, and 2.7 if it is included.

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4.7 Uncertainties and assessments of results

The estimation of NO_x emissions is based on the use of several quantities with varying uncertainties. The error bars shown in Figures 5 and 9 are thus calculated from uncertainty statistics whose references are presented in this section. Since these references do not specify the exact nature of these statistics, we assume they correspond to standard deviations. The uncertainty of tropospheric NO_2 columns under polluted conditions is dominated by the sensitivity of satellite observations to lower tropospheric air masses, expressed by the tropospheric air-mass factor (AMF). The column relative uncertainty due to the AMF is of the order of 30% (Boersma et al., 2004 [43]). S-5P validation activities indicate that TROPOMI tropospheric NO_2 columns are systematically biased low by about 30%-50% over cities (Compernolle et al., 2018 [81]), which is most likely related to the *a priori* profiles used within the operational retrieval that do not reflect well the NO_2 peak close to ground. For the Middle East region, the impact of the *a priori* profile is less critical, as surface albedo is generally high and cloud fractions are generally low. Thus, we expect no such bias, and consider a relative uncertainty of 30% for the tropospheric column. Other uncertainties must be taken into account: the transition from NO_2 TROPOMI columns to NO_x emissions requires parameters which appear in Equation (2) and Equation (3). For wind module, uncertainties are generally of about 1 m/s for components taken at precise altitudes (Coburn et al., 2019 [82]; Beirle et al., 2019 [15]). Here, we assume an uncertainty of 3 m/s for both zonal and meridional wind components. For [OH], the analysis of different methods conducted by Huijnen et al., 2019 [83] showed smaller differences for low latitudes than for extratropics, but still significant. We thus take a relative uncertainty of 30% for OH concentration. For the reaction rate k_{mean}, the value of the corresponding relative uncertainty has been estimated by Burkholder et al., 2020 [30]. Finally, we use the sensitivity tests performed in Section 4.3 to assess the uncertainty associated with the choice of the vertical level. The cumulative effects on the final emissions of the three parameters studied, in particular the OH concentration, lead to a relative uncertainty that varies from month to month between 7 and 18%. The propagation of these different uncertainties on the monthly estimates of NO_x emissions in Egypt leads to an expanded uncertainty between 47 and 51%. For lifetimes calculated with the EMG function fitting, the corresponding expanded uncertainty ranges between 18% and 79%.

We acknowledge the fact that our treatment of NO_x is simplified. Many minor sinks highlighted in Section 3.1 are not taken into account. In particular, anthropogenic VOC emissions, which remove NO_x from the atmosphere, compete with the oxidation by OH for the representation of NO_x loss. These emissions are difficult to estimate and the corresponding sink is complex to model. Taking this reaction into account would *a priori* lead to a strengthening of the sink term and thus to an increase of the NO_x emissions estimates. Other assumptions in the model are also simplifications. For instance, obtaining anthropogenic emissions by subtracting the average emissions over background cells assumes that the non-anthropogenic sources of NO_2 are similar inside and outside the mask, which is not true, since a large part of the mask cells correspond to croplands. For these cells, soil emissions may play a non-negligible role in the natural NO_2 budget. As a consequence, mean background emissions that are removed from NO_x emissions estimates above mask cells might be under-estimated. Finally, the reliability of the data used can be questioned. The representation of the wind is crucial to avoid creating artificial patterns in the transport term. The OH concentration, which is proportional to the intensity of the sink term, is also important. We have shown that OH concentrations are partially responsible for an important drop in NO_x emissions in the winter of 2019/2020 that may be unrealistic. Because this decrease is largely due to variations in OH concentrations provided by CAMS, whose reliability has been evaluated for Riyadh, then the transposability hypothesis between Riyadh and Egypt may be subject to further discussion.

5   Conclusions

In this study, we investigated the potential of a top-down model of NO_x emissions based on TROPOMI retrievals at high resolution over Egypt. The model is based on the study of a transport term and a sink term that requires different parameters to be calculated. Among those parameters, the concentration in OH, involved in the calculation of the NO_2 mixed lifetime, is of fundamental importance. The comparison between NO_2 lifetimes derived from OH concentrations and NO_2 lifetimes derived from EMG function fittings of line density profiles shows that the OH concentration provided by CAMS is reasonably reliable for the country. Parameters are therefore taken in the first 200 m of the planetary boundary layer, because it is where OH shows the best consistency. However, the vertical sensitivity linked to this parameters remains high. Results illustrate the importance of the transport term at local scale, which is of the same order of magnitude as the sink term above large cities and industrial facilities; it ceases to be relevant only at the scale of the whole country. The top-down model is able to characterise declines in human activities due to restrictions during the COVID-19 pandemic or to Friday rest. It also estimates higher emissions during summer. These high emissions may be interpreted by a higher consumption of electricity driven by air-conditioning during hot days, but it

remains unclear whether this pattern clearly reproduces changes in human activity, in particular because the emission inventories show different seasonalities. These inventories also differ in the amount of total emissions: the average value for TROPOMI-derived NO_x emissions is 7.0% higher than CAMS estimates. This discrepancy could be resolved by comparing the results of the model and inventory estimates to industrial production or electricity consumption data at the scale of countries or regions.

This study demonstrates the potential of TROPOMI data for evaluating NO_x emissions in the Middle East region. More generally, it demonstrates the importance of the contribution of independent observation systems to overcome the weaknesses of emission inventories, provided that the local chemistry is well understood and modelled. The development of similar applications for different species is likely to allow better monitoring of global anthropogenic emissions, therefore helping companies and countries to report their emissions of air pollutants and greenhouse gases as part of their strategies and obligations to tackle air pollution issues and climate change.

References

[1] A. Baklanov, L. T. Molina, and M. Gauss, "Megacities, air quality and climate," *Atmospheric Environment*, vol. 126, pp. 235–249, 2016.

[2] L. Tang, X. Xue, M. Jia, H. Jing, T. Wang, R. Zhen, M. Huang, J. Tian, J. Guo, L. Li, *et al.*, "Iron and steel industry emissions and contribution to the air quality in China," *Atmospheric Environment*, vol. 237, p. 117668, 2020.

[3] H. C. Kim, C. Bae, M. Bae, O. Kim, B.-U. Kim, C. Yoo, J. Park, J. Choi, J.-b. Lee, B. Lefer, *et al.*, "Space-Borne monitoring of NOx emissions from cement kilns in South Korea," *Atmosphere*, vol. 11, no. 8, p. 881, 2020.

[4] A. Singh and M. Agrawal, "Acid rain and its ecological consequences," *Journal of Environmental Biology*, vol. 29, no. 1, p. 15, 2007.

[5] U. EPA, "Integrated science assessment for oxides of nitrogen–health criteria," *US Environmental Protection Agency, Washington, DC [Google Scholar]*, 2016.

[6] M. Crippa, G. Janssens-Maenhout, F. Dentener, D. Guizzardi, K. Sindelarova, M. Muntean, R. Van Dingenen, and C. Granier, "Forty years of improvements in European air quality: regional policy-industry interactions with global impacts," *Atmospheric Chemistry and Physics*, vol. 16, no. 6, pp. 3825–3841, 2016.

[7] A. El-Magd, N. Zanaty, E. M. Ali, H. Irie, A. I. Abdelkader, *et al.*, "Investigation of aerosol climatology, optical characteristics and variability over Egypt based on satellite observations and in-situ measurements," *Atmosphere*, vol. 11, no. 7, p. 714, 2020.

[8] UNEP (United Nations Environment Programme), "Air quality policies in Egypt," 2015.

[9] B. Xue and W. Ren, "China's uncertain CO2 emissions," *Nature Climate Change*, vol. 2, no. 11, pp. 762–762, 2012.

[10] C. Leue, M. Wenig, T. Wagner, O. Klimm, U. Platt, and B. Jähne, "Quantitative analysis of NOx emissions from global ozone Monitoring Experiment satellite image sequences," *Journal of Geophysical Research: Atmospheres*, vol. 106, no. D6, pp. 5493–5505, 2001.

[11] R. V. Martin, D. J. Jacob, K. Chance, T. P. Kurosu, P. I. Palmer, and M. J. Evans, "Global inventory of nitrogen oxide emissions constrained by space-based observations of NO2 columns," *Journal of Geophysical Research: Atmospheres*, vol. 108, no. D17, 2003.

[12] B. Mijling and R. Van Der A, "Using daily satellite observations to estimate emissions of short-lived air pollutants on a mesoscopic scale," *Journal of Geophysical Research: Atmospheres*, vol. 117, no. D17, 2012.

[13] B. de Foy, Z. Lu, D. G. Streets, L. N. Lamsal, and B. N. Duncan, "Estimates of power plant NOx emissions and lifetimes from OMI NO2 satellite retrievals," *Atmospheric Environment*, vol. 116, pp. 1–11, 2015.

[14] D. L. Goldberg, Z. Lu, D. G. Streets, B. de Foy, D. Griffin, C. A. McLinden, L. N. Lamsal, N. A. Krotkov, and H. Eskes, "Enhanced Capabilities of TROPOMI NO2: Estimating NOx from North American Cities and Power Plants," *Environmental science & technology*, vol. 53, no. 21, pp. 12594–12601, 2019.

[15] S. Beirle, C. Borger, S. Dörner, A. Li, Z. Hu, F. Liu, Y. Wang, and T. Wagner, "Pinpointing nitrogen oxide emissions from space," *Science advances*, vol. 5, no. 11, p. eaax9800, 2019.

[16] A. Lorente, K. Boersma, H. Eskes, J. Veefkind, J. Van Geffen, M. De Zeeuw, H. D. van der Gon, S. Beirle, and M. Krol, "Quantification of nitrogen oxides emissions from build-up of pollution over Paris with TROPOMI," *Scientific reports*, vol. 9, no. 1, pp. 1–10, 2019.

[17] K. Lange, A. Richter, and J. P. Burrows, "Variability of nitrogen oxide emission fluxes and lifetimes estimated from Sentinel-5P TROPOMI observations," *Atmospheric Chemistry and Physics Discussions*, pp. 1–32, 2021.

[18] J. Veefkind, I. Aben, K. McMullan, H. Förster, J. De Vries, G. Otter, J. Claas, H. Eskes, J. De Haan, Q. Kleipool, *et al.*, "TROPOMI on the ESA Sentinel-5 Precursor: A GMES mission for global observations of the atmospheric composition for climate, air quality and ozone layer applications," *Remote sensing of environment*, vol. 120, pp. 70–83, 2012.

[19] S. Lama, S. Houweling, K. F. Boersma, H. Eskes, I. Aben, H. A. Denier van der Gon, M. C. Krol, H. Dolman, T. Borsdorff, and A. Lorente, "Quantifying burning efficiency in megacities using the NO2/CO ratio from the Tropospheric Monitoring Instrument (TROPOMI)," *Atmospheric Chemistry and Physics*, vol. 20, no. 17, pp. 10295–10310, 2020.

[20] H. Eskes, K. Eichmann, J. Lambert, D. Loyola, J. Veefkind, A. Dehn, and C. Zehner, "S5P Mission Performance Centre Nitrogen Dioxide [L2_NO2] readme," *Royal Netherlands Meteorological Institute (KNMI) De Bilt, the Netherlands, version*, vol. 1, no. 00, 2019.

[21] H. Hersbach, B. Bell, P. Berrisford, S. Hirahara, A. Horányi, J. Muñoz-Sabater, J. Nicolas, C. Peubey, R. Radu, D. Schepers, *et al.*, "The ERA5 global reanalysis," *Quarterly Journal of the Royal Meteorological Society*, vol. 146, no. 730, pp. 1999–2049, 2020.

[22] V. Huijnen, H. Eskes, A. Wagner, M. Schulz, Y. Christophe, M. Ramonet, S. Basart, A. Benedictow, A.-M. Blechschmidt, S. Chabrillat, *et al.*, "Validation report of the CAMS near-real-time global atmospheric composition service: System evolution and performance statistics. Status up to 1 June 2016; https://pure.mpg.de/rest/items/item_2441827/component/file_2441834/content," 2016.

[23] J. H. Seinfeld, "Urban air pollution: state of the science," *Science*, vol. 243, no. 4892, pp. 745–752, 1989.

[24] J. A. Logan, M. J. Prather, S. C. Wofsy, and M. B. McElroy, "Tropospheric chemistry: a global perspective," *Journal of Geophysical Research: Oceans*, vol. 86, no. C8, pp. 7210–7254, 1981.

[25] H. Levy, "Normal atmosphere: Large radical and formaldehyde concentrations predicted," *Science*, vol. 173, no. 3992, pp. 141–143, 1971.

[26] L. Valin, A. Russell, R. Hudman, and R. Cohen, "Effects of model resolution on the interpretation of satellite NO2 observations," *Atmospheric Chemistry and Physics*, vol. 11, no. 22, pp. 11647–11655, 2011.

[27] J. Lelieveld, S. Gromov, A. Pozzer, and D. Taraborrelli, "Global tropospheric hydroxyl distribution, budget and reactivity," *Atmospheric Chemistry and Physics*, vol. 16, no. 19, pp. 12477–12493, 2016.

[28] M. Li, E. Karu, C. Brenninkmeijer, H. Fischer, J. Lelieveld, and J. Williams, "Tropospheric OH and stratospheric OH and Cl concentrations determined from CH4, CH3Cl, and SF6 measurements," *NPJ Climate and Atmospheric Science*, vol. 1, no. 1, pp. 1–7, 2018.

[29] G. M. Wolfe, J. M. Nicely, J. M. S. Clair, T. F. Hanisco, J. Liao, L. D. Oman, W. B. Brune, D. Miller, A. Thames, G. G. Abad, *et al.*, "Mapping hydroxyl variability throughout the global remote troposphere via synthesis of airborne and satellite formaldehyde observations," *Proceedings of the National Academy of Sciences*, vol. 116, no. 23, pp. 11171–11180, 2019.

[30] J. Burkholder, S. Sander, J. Abbatt, J. Barker, C. Cappa, J. Crounse, T. Dibble, R. Huie, C. Kolb, M. Kurylo, *et al.*, "Chemical kinetics and photochemical data for use in atmospheric studies; evaluation number 19," tech. rep., Pasadena, CA: Jet Propulsion Laboratory, National Aeronautics and Space . . . , 2020.

[31] J. Yienger and H. Levy, "Empirical model of global soil-biogenic NOx emissions," *Journal of Geophysical Research: Atmospheres*, vol. 100, no. D6, pp. 11447–11464, 1995.

[32] J. J. Hoelzemann, M. G. Schultz, G. P. Brasseur, C. Granier, and M. Simon, "Global wildland fire emission model (GWEM): Evaluating the use of global area burnt satellite data," *Journal of Geophysical Research: Atmospheres*, vol. 109, no. D14, 2004.

[33] D. H. Ehhalt, F. Rohrer, and A. Wahner, "Sources and distribution of NOx in the upper troposphere at northern mid-latitudes," *Journal of Geophysical Research: Atmospheres*, vol. 97, no. D4, pp. 3725–3738, 1992.

[34] L. Jaeglé, L. Steinberger, R. V. Martin, and K. Chance, "Global partitioning of NOx sources using satellite observations: Relative roles of fossil fuel combustion, biomass burning and soil emissions," *Faraday discussions*, vol. 130, pp. 407–423, 2005.

[35] J.-F. Müller and T. Stavrakou, "Inversion of CO and NOx emissions using the adjoint of the images model," *Atmospheric Chemistry and Physics*, vol. 5, no. 5, pp. 1157–1186, 2005.

[36] J.-T. Lin, "Satellite constraint for emissions of nitrogen oxides from anthropogenic, lightning and soil sources over East China on a high-resolution grid," *Atmospheric Chemistry and Physics*, vol. 12, no. 6, pp. 2881–2898, 2012.

[37] CIESIN, "CIESIN. Global Rural-Urban Mapping Project, Version 1 (GRUMPv1). Center for International Earth Science Information Network - CIESIN - Columbia University, International Food Policy Research Institute - IFPRI, The World Bank, and Centro Internacional de Agricultura Tropical - CIAT. 2011, Palisades, NY: NASA Socioeconomic Data and Applications Center (SEDAC)," 2019.

[38] C. D. Elvidge, M. Zhizhin, K. Baugh, F.-C. Hsu, and T. Ghosh, "Methods for global survey of natural gas flaring from visible infrared imaging radiometer suite data," *Energies*, vol. 9, no. 1, p. 14, 2016.

[39] M. Crippa, D. Guizzardi, M. Muntean, E. Schaaf, E. Solazzo, F. Monforti-Ferrario, J. Olivier, and E. Vignati, "Fossil CO2 emissions of all world countries," *Luxembourg: European Commission*, pp. 1–244, 2020.

[40] C. Granier, S. Darras, H. D. van der Gon, D. Jana, N. Elguindi, G. Bo, G. Michael, G. Marc, J.-P. Jalkanen, J. Kuenen, *et al.*, *The Copernicus atmosphere monitoring service global and regional emissions (April 2019 version)*. PhD thesis, Copernicus Atmosphere Monitoring Service, 2019.

[41] R. M. Hoesly, S. J. Smith, L. Feng, Z. Klimont, G. Janssens-Maenhout, T. Pitkanen, J. J. Seibert, L. Vu, R. J. Andres, R. M. Bolt, *et al.*, "Historical (1750–2014) anthropogenic emissions of reactive gases and aerosols from the community emissions data system (CEDS)," *Geoscientific Model Development*, vol. 11, no. 1, pp. 369–408, 2018.

[42] G. Janssens-Maenhout, M. Crippa, D. Guizzardi, M. Muntean, E. Schaaf, F. Dentener, P. Bergamaschi, V. Pagliari, J. G. Olivier, J. A. Peters, *et al.*, "Edgar v4. 3.2 global atlas of the three major greenhouse gas emissions for the period 1970–2012," *Earth System Science Data*, vol. 11, no. 3, pp. 959–1002, 2019.

[43] K. Boersma, H. Eskes, and E. Brinksma, "Error analysis for tropospheric NO2 retrieval from space," *Journal of Geophysical Research: Atmospheres*, vol. 109, no. D4, 2004.

[44] A. Lorente, K. Folkert Boersma, H. Yu, S. Dörner, A. Hilboll, A. Richter, M. Liu, L. N. Lamsal, M. Barkley, I. D. Smedt, *et al.*, "Structural uncertainty in air mass factor calculation for NO2 and HCHO satellite retrievals," *Atmospheric Measurement Techniques*, vol. 10, no. 3, pp. 759–782, 2017.

[45] S. Sander, R. Friedl, J. Abbatt, J. Barker, J. Burkholder, D. Golden, C. Kolb, M. Kurylo, G. Moortgat, P. Wine, *et al.*, "Chemical kinetics and photochemical data for use in atmospheric studies, jpl publication 10-6," *Evaluation*, no. 17, 2011.

[46] B. A. Nault, C. Garland, P. J. Wooldridge, W. H. Brune, P. Campuzano-Jost, J. D. Crounse, D. A. Day, J. Dibb, S. R. Hall, L. G. Huey, *et al.*, "Observational constraints on the oxidation of NOx in the upper troposphere," *The Journal of Physical Chemistry A*, vol. 120, no. 9, pp. 1468–1478, 2016.

[47] T. Stavrakou, J.-F. Müller, K. Boersma, R. Van Der A, J. Kurokawa, T. Ohara, and Q. Zhang, "Key chemical NOx sink uncertainties and how they influence top-down emissions of nitrogen oxides," *Atmospheric Chemistry and Physics*, vol. 13, no. 17, pp. 9057–9082, 2013.

[48] H. Fang, F. Baret, S. Plummer, and G. Schaepman-Strub, "An overview of global leaf area index (LAI): Methods, products, validation, and applications," *Reviews of Geophysics*, vol. 57, no. 3, pp. 739–799, 2019.

[49] E. R. Delaria, B. K. Place, A. X. Liu, and R. C. Cohen, "Laboratory measurements of stomatal NO2 deposition to native california trees and the role of forests in the NOx cycle," *Atmospheric Chemistry and Physics*, vol. 20, no. 22, pp. 14023–14041, 2020.

[50] N. Sobanski, J. Thieser, J. Schuladen, C. Sauvage, W. Song, J. Williams, J. Lelieveld, and J. N. Crowley, "Day and night-time formation of organic nitrates at a forested mountain site in south-west germany," *Atmospheric Chemistry and Physics*, vol. 17, no. 6, pp. 4115–4130, 2017.

[51] A. Guenther, T. Karl, P. Harley, C. Wiedinmyer, P. I. Palmer, and C. Geron, "Estimates of global terrestrial isoprene emissions using megan (model of emissions of gases and aerosols from nature)," *Atmospheric Chemistry and Physics*, vol. 6, no. 11, pp. 3181–3210, 2006.

[52] P. S. Romer Present, A. Zare, and R. C. Cohen, "The changing role of organic nitrates in the removal and transport of NOx," *Atmospheric Chemistry and Physics*, vol. 20, no. 1, pp. 267–279, 2020.

[53] N. Butkovskaya, A. Kukui, N. Pouvesle, and G. Le Bras, "Formation of nitric acid in the gas-phase HO2+NO reaction: Effects of temperature and water vapor," *The Journal of Physical Chemistry A*, vol. 109, no. 29, pp. 6509–6520, 2005.

[54] N. Butkovskaya, M.-T. Rayez, J.-C. Rayez, A. Kukui, and G. Le Bras, "Water vapor effect on the HNO3 yield in the HO2+NO reaction: experimental and theoretical evidence," *The Journal of Physical Chemistry A*, vol. 113, no. 42, pp. 11327–11342, 2009.

[55] N. Friedrich, P. Eger, J. Shenolikar, N. Sobanski, J. Schuladen, D. Dienhart, B. Hottmann, I. Tadic, H. Fischer, M. Martinez, *et al.*, "Reactive nitrogen around the arabian peninsula and in the mediterranean sea during the 2017 AQABA ship campaign," *Atmospheric Chemistry and Physics*, vol. 21, no. 10, pp. 7473–7498, 2021.

[56] W. Moxim, H. Levy, and P. Kasibhatla, "Simulated global tropospheric PAN: Its transport and impact on NOx," *Journal of Geophysical Research: Atmospheres*, vol. 101, no. D7, pp. 12621–12638, 1996.

[57] E. Fischer, D. J. Jacob, R. M. Yantosca, M. P. Sulprizio, D. Millet, J. Mao, F. Paulot, H. Singh, A. Roiger, L. Ries, *et al.*, "Atmospheric peroxyacetyl nitrate (PAN): a global budget and source attribution," *Atmospheric Chemistry and Physics*, vol. 14, no. 5, pp. 2679–2698, 2014.

[58] C. A. Longfellow, A. Ravishankara, and D. R. Hanson, "Reactive uptake on hydrocarbon soot: Focus on NO2," *Journal of Geophysical Research: Atmospheres*, vol. 104, no. D11, pp. 13833–13840, 1999.

[59] S. Attia, A. Evrard, and E. Gratia, "Development of benchmark models for the Egyptian residential buildings sector," *Applied Energy*, vol. 94, pp. 270–284, 2012.

[60] EEHC, "Egyptian Electricity Holding Company annual report 2019/2020," 2021.

[61] F. Rohrer and H. Berresheim, "Strong correlation between levels of tropospheric hydroxyl radicals and solar ultraviolet radiation," *Nature*, vol. 442, no. 7099, pp. 184–187, 2006.

[62] S. Beirle, K. F. Boersma, U. Platt, M. G. Lawrence, and T. Wagner, "Megacity emissions and lifetimes of nitrogen oxides probed from space," *Science*, vol. 333, no. 6050, pp. 1737–1739, 2011.

[63] L. Valin, A. Russell, and R. Cohen, "Variations of OH radical in an urban plume inferred from NO2 column measurements," *Geophysical Research Letters*, vol. 40, no. 9, pp. 1856–1860, 2013.

[64] M. Filioglou, E. Giannakaki, J. Backman, J. Kesti, A. Hirsikko, R. Engelmann, E. O'Connor, J. T. Leskinen, X. Shang, H. Korhonen, *et al.*, "Optical and geometrical aerosol particle properties over the United Arab Emirates," *Atmospheric Chemistry and Physics*, vol. 20, no. 14, pp. 8909–8922, 2020.

[65] M. Khoder, "Diurnal, seasonal and weekdays–weekends variations of ground level ozone concentrations in an urban area in greater cairo," *Environmental Monitoring and Assessment*, vol. 149, no. 1, pp. 349–362, 2009.

[66] T. Graedel, L. Farrow, and T. Weber, "Kinetic studies of the photochemistry of the urban troposphere," *Atmospheric Environment (1967)*, vol. 10, no. 12, pp. 1095–1116, 1976.

[67] J. H. Seinfeld and S. N. Pandis, "Atmospheric chemistry and physics from air pollution to climate change," 2006.

[68] S. Beirle, C. Borger, S. Dörner, H. Eskes, V. Kumar, A. de Laat, and T. Wagner, "Catalog of NOx emissions from point sources as derived from the divergence of the NO2 flux for TROPOMI," *Earth System Science Data*, vol. 13, no. 6, pp. 2995–3012, 2021.

[69] T. Stavrakou, J.-F. Müller, M. Bauwens, K. Boersma, and J. van Geffen, "Satellite evidence for changes in the NO2 weekly cycle over large cities," *Scientific reports*, vol. 10, no. 1, pp. 1–9, 2020.

[70] U. IEA, "Global energy review 2020," *Ukraine.[Online] https://www. iea. org/countries/ukraine [Accessed: 2020-09-10]*, 2020.

[71] Z. S. Venter, K. Aunan, S. Chowdhury, and J. Lelieveld, "COVID-19 lockdowns cause global air pollution declines," *Proceedings of the National Academy of Sciences*, vol. 117, no. 32, pp. 18984–18990, 2020.

[72] M. Bauwens, S. Compernolle, T. Stavrakou, J.-F. Müller, J. Van Gent, H. Eskes, P. F. Levelt, R. van der A, J. Veefkind, J. Vlietinck, *et al.*, "Impact of coronavirus outbreak on NO2 pollution assessed using TROPOMI and OMI observations," *Geophysical Research Letters*, vol. 47, no. 11, p. e2020GL087978, 2020.

[73] G. I. Gkatzelis, J. B. Gilman, S. S. Brown, H. Eskes, A. R. Gomes, A. C. Lange, B. C. McDonald, J. Peischl, A. Petzold, C. R. Thompson, *et al.*, "The global impacts of COVID-19 lockdowns on urban air pollutiona critical review and recommendations," *Elementa: Science of the Anthropocene*, vol. 9, no. 1, 2021.

[74] T. Hale, N. Angrist, R. Goldszmidt, B. Kira, A. Petherick, T. Phillips, S. Webster, E. Cameron-Blake, L. Hallas, S. Majumdar, *et al.*, "A global panel database of pandemic policies (oxford covid-19 government response tracker)," *Nature Human Behaviour*, vol. 5, no. 4, pp. 529–538, 2021.

[75] M. M. El-Sheekh and I. A. Hassan, "Lockdowns and reduction of economic activities during the COVID-19 pandemic improved air quality in Alexandria, Egypt," *Environmental Monitoring and Assessment*, vol. 193, no. 1, pp. 1–7, 2021.

[76] I. Abou El-Magd and N. Zanaty, "Impacts of short-term lockdown during COVID-19 on air quality in Egypt," *The Egyptian Journal of Remote Sensing and Space Science*, 2020.

[77] L. Abdallah and T. El-Shennawy, "Evaluation of CO2 emission from Egypt's future power plants," *Euro-Mediterranean Journal for Environmental Integration*, vol. 5, no. 3, pp. 1–8, 2020.

[78] A. M. A. H. Elharidi, P. G. Tuohy, and M. Teamah, "Facing the growing problem of the electric power consumption in Egyptian residential building using building performance simulation program," in *Building simulation Cairo 2013 conference*, 2013.

[79] M. M. Nassief, "Evaluation of electricity consumption of a residential flat in egypt," *American Journal of Electrical Power and Energy Systems*, vol. 3, no. 2, pp. 7–44, 2014.

[80] S. M. Wahba, B. A. Kamel, K. M. Nassar, and A. S. Abdelsalam, "Effectiveness of green roofs and green walls on energy consumption and indoor comfort in arid climates," *Civil Engineering Journal*, vol. 4, no. 10, pp. 2284–2295, 2018.

[81] S. Compernolle *et al.*, "S5P MPC VDAF validation web article: Nitrogen dioxide," tech. rep., S5P-MPC-VDAF-WVA-L2_NO2_20180904, 2018.

[82] J. J. Coburn, "Assessing wind data from reanalyses for the upper midwest," *Journal of Applied Meteorology and Climatology*, vol. 58, no. 3, pp. 429–446, 2019.

[83] V. Huijnen, A. Pozzer, J. Arteta, G. Brasseur, I. Bouarar, S. Chabrillat, Y. Christophe, T. Doumbia, J. Flemming, J. Guth, *et al.*, "Quantifying uncertainties due to chemistry modelling–evaluation of tropospheric composition simulations in the CAMS model (cycle 43r1)," *Geoscientific Model Development*, vol. 12, no. 4, pp. 1725–1752, 2019.