



Heterogeneous Interactions between SO₂ and Organic Peroxides in Submicron Aerosol

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1 **Abstract**

2 Atmospheric models often underestimate particulate sulfate, a major component in ambient
3 aerosol, suggesting missing sulfate formation mechanisms in the models. Heterogeneous
4 reactions between SO₂ and aerosol play an important role in particulate sulfate formation and its
5 physicochemical evolution. Here we study the reactive uptake kinetics of SO₂ onto aerosol
6 containing organic peroxides. We present chamber studies of SO₂ reactive uptake performed
7 under different relative humidities (RH), particulate peroxide contents, peroxide types, and
8 aerosol acidities. Using different model organic peroxides mixed with ammonium sulfate
9 particles, SO₂ uptake coefficient (γ_{SO_2}) was found to be exponentially dependent on RH. γ_{SO_2}
10 increases from 10⁻³ at RH 25% to 10⁻² at RH 71% as measured for a multifunctional organic
11 peroxide. Under similar conditions, the kinetics were found to be structurally dependent:
12 multifunctional organic peroxides have a higher γ_{SO_2} than those with only one peroxide group,
13 consistent with the reactivity trend observed previously in the aqueous phase. In addition, γ_{SO_2} is
14 linearly related to particle-phase peroxide content, which in turn depends on gas-particle
15 partitioning of organic peroxides. Aerosol acidity plays a complex role in determining SO₂
16 uptake rate, influenced by the effective Henry's Law constant of SO₂ and the condensed phase
17 kinetics of the peroxide-SO₂ reaction in the highly concentrated aerosol phase. These uptake
18 coefficients are consistently higher than those calculated from the reaction kinetics in the bulk
19 aqueous phase, and we show experimental evidence suggesting that other factors, such as
20 particle-phase ionic strength, can play an essential role in determining the uptake kinetics. γ_{SO_2}
21 for different types of secondary organic aerosol (SOA) were measured to be on the order of 10⁻⁴.
22 Overall, this study provides quantitative evidence of the multiphase reactions between SO₂ and
23 organic peroxides, highlighting the important factors that govern the uptake kinetics.



24 **Introduction**

25 Sulfate and organic compounds are ubiquitous particulate components in both polluted and
26 pristine environments (Chen et al., 2009;Andreae et al., 2018;He et al., 2011;Sun et al.,
27 2013;Huang et al., 2014), with important implications for public health and global climate
28 (Hallquist et al., 2009). Particulate sulfate can form via S(IV) oxidation by OH radicals in the gas
29 phase and via oxidation in cloud water, fog droplets or the aerosol aqueous phase, including by
30 H₂O₂, O₂ (catalyzed by transition metals), O₃, NO₂ and small organic peroxides (methyl
31 hydroperoxide and peroxyacetic acid) (Seinfeld and Pandis, 2012). However, atmospheric
32 models tend to underestimate particulate sulfate production on both global (Tie et al., 2001;Yang
33 et al., 2017;Fairlie et al., 2010) and regional scales, especially during heavy haze episodes (Wang
34 et al., 2014;Zheng et al., 2015;Sha et al., 2019;Gao et al., 2016;Li et al., 2017;Huang et al.,
35 2019), suggesting that the overall kinetics may be underestimated and/or important mechanisms
36 may be missing in models.

37 To reconcile these differences, studies have investigated novel reaction mechanisms of sulfate
38 formation. Stabilized Criegee intermediates (sCIs) were found to oxidize SO₂ rapidly and
39 proposed to be an important source of ambient sulfate (Mauldin et al., 2012). When Liu et al.
40 (2019) applied this mechanism and kinetics to a source-oriented WRF-Chem model, the sCIs
41 pathway was found to only account for at most 9% of the total particulate sulfate. Reactive
42 nitrogen species (such as NO₂) have also been put forward to account for the missing sulfate at
43 relatively high aerosol pH (close to 7) (Wang et al., 2016;Cheng et al., 2016). However, such
44 high aerosol pH is not substantiated by thermodynamic models, which conclude that pH ranges
45 between 4 and 5 even in polluted regions (Song et al., 2018;Guo et al., 2017). A recent modeling
46 study incorporating this heterogeneous NO_x mechanism still exhibited a discrepancy of 20%



47 between the predicted and observed sulfate, indicating the possibility of unknown mechanisms
48 (Huang et al., 2019). Other factors may play a role in enhancing the particle-phase sulfate
49 formation rates. Chen et al. (2019) investigated the synergistic effects of NO₂ and NH₃ on sulfate
50 formation, and found that the rate of this reaction can be enhanced by the high ionic strength in
51 the particle phase. This enhancement effect by solute strength on sulfate formation was also
52 investigated for the H₂O₂ pathway in aerosol liquid water. Liu et al. (2020) found ionic strength
53 and general acid-catalyzed mechanisms can cause the S(VI) formation rate to be nearly 50 times
54 faster in aerosol phase than in dilute solutions. On the other hand, during the severe haze
55 episodes in China (Li et al., 2020; Guo et al., 2017), transition metal ion (TMI) catalysis of SO₂
56 oxidation by O₂ can be significantly suppressed in the aerosol phase due to high ionic strength
57 (Liu et al., 2020; Cheng et al., 2016; Su et al., 2020).

58 In addition to high solute strength, submicron aerosol is also rich in organic compounds (Jimenez
59 et al., 2009; Hallquist et al., 2009). In recent years, many studies have investigated the potential
60 role of heterogeneous interactions between SO₂ and organic aerosol on particulate sulfate
61 formation. Song et al. (2019) found heterogeneous oxidation of hydroxymethanesulfonate
62 (HMS) by OH can trigger rapid sulfate formation. Wang et al. (2020) studied photosensitizers in
63 ambient particles and found this pathway could be essential under specific light conditions.

64 Recent studies found reactive intermediates from isoprene oxidation (Huang et al., 2019) and
65 benzoic acid (Huang et al., 2020), can yield a variety of organosulfur species upon catalysis by
66 TMI. Other studies have also investigated the interactions between secondary organic aerosol
67 (SOA) and SO₂. Field observations found that ambient sulfate abundance is highly correlated
68 with SOA formation (Yee et al., 2020; Xu et al., 2015). Liu et al. (2019) found that SO₂ enhances
69 SOA formation and average carbon oxidation state during methoxyphenol photooxidation. By



70 performing chamber experiments with limonene SOA formation in the presence of SO₂, Ye et al.
71 (2018) also observed significant SO₂ decay along with increased SOA yields and carbon
72 oxidation state, proposing that organic peroxides in SOA may be the key reactive intermediates
73 for SO₂ oxidation.

74 Organic peroxides are key intermediates for aerosol formation and ubiquitously exist in many
75 SOA systems (Hallquist et al., 2009; Bianchi et al., 2019). Numerous studies have reported
76 peroxide content of 20–60% for isoprene and monoterpene derived SOA (Surratt et al., 2006; Ng
77 et al., 2008; Ye et al., 2018; Epstein et al., 2014). A significant fraction of organic peroxide (30%–
78 50%) has also been found in naphthalene-derived SOA under low/high NO_x conditions
79 (Kautzman et al., 2009). Using model simulations, Bonn et al. (2004) found organic
80 hydroperoxides can account for up to 60% of global SOA. The aqueous phase reaction kinetics
81 between organic peroxides and dissolved SO₂ have been explored in previous studies (Lind et al.,
82 1987; Gunz and Hoffmann, 1990; Wang et al., 2019; Dovrou et al., 2019; Yao et al., 2019). The
83 second order reaction rate constants for organic peroxides in SOA (Dovrou et al., 2019; Yao et
84 al., 2019) and S(IV) were measured to be on the order of 10²–10³ M⁻¹ s⁻¹, which are within the
85 range of those measured for commercially available organic peroxides (Wang et al., 2019) and
86 small organic peroxides (Lind et al., 1987). Yao et al. (2019) quantified the reactive uptake
87 coefficient of SO₂ (γ_{SO_2}) onto α -pinene SOA to be on the order of 10⁻⁴–10⁻³, which is positively
88 dependent on RH and inferred particle-phase peroxide content. These reactions are also linked to
89 the formation of organosulfates (Wang et al., 2019). Both inorganic sulfate (85–90%) and
90 organosulfates (10–15%) were observed as products of SO₂ reactive uptake onto SOA (Yao et al.,
91 2019).



92 Given the potential significance of SO₂ reactive uptake in particulate sulfate formation, a more
93 in-depth study is needed to determine the important factors that govern the heterogeneous
94 kinetics of SO₂ onto organic peroxide containing aerosol. In this study, we measured γ_{SO_2} for two
95 categories of aerosol: 1. Model organic peroxides mixed with ammonium sulfate or malonic acid
96 and 2. SOA from a few representative biogenic and anthropogenic precursors. The impacts of
97 RH, peroxide type, peroxide content, and condensed phase pH on SO₂ reactive uptake were
98 evaluated systematically with the goal of better understanding atmospheric multiphase sulfate
99 formation.

100

101 **2. Methods**

102 The reactive uptake of SO₂ onto peroxide-containing particles was studied in a 1 m³ Teflon
103 chamber under ambient temperature and pressure. In brief, generated particles and SO₂ were
104 introduced into the chamber separately. The consumption of SO₂, changes in particle size
105 distribution and chemical composition were monitored to estimate the reactive uptake
106 coefficients. Particles were also collected on filters for offline chemical characterization.

107

108 **2.1 Seed aerosol generation**

109 In this work, two types of aerosol were used to investigate the uptake of SO₂. The first is
110 ammonium sulfate or malonic acid mixed with model organic peroxides (Fig. S1). In this first set
111 of experiments, an aerosol atomizer (Model 3076, TSI Inc., USA) was used to generate aqueous
112 particles from dilute solution. Each solution consists of ammonium sulfate ($\geq 99\%$, Sigma-
113 Aldrich) or malonic acid (99%, Sigma-Aldrich) and a model organic peroxide in ultrapure water
114 (HPLC grade, Fisher Chemical). For the experiments investigating the relationship between γ_{SO_2}



115 and peroxide type (Expt. 2-14), different commercially available organic peroxides were used,
116 including tert-butyl hydroperoxide (70 wt. % in water, Sigma-Aldrich), cumene hydroperoxide
117 (80 wt. % in water, Sigma-Aldrich), and 2-butanone peroxide (40% wt. % in water, Sigma-
118 Aldrich). The molar ratio of organic peroxide to ammonium sulfate in the atomizing solution was
119 2:1 with the aim of being atmospherically relevant (corresponding to maximum particulate
120 peroxide molar fraction of 66% and mass fraction of approximately 50-70% if all the organic
121 peroxides were assumed to remain in the particle phase). This ratio was used as a proxy for total
122 peroxide content in both gas and particle phase relative to that of ammonium sulfate upon
123 atomization. For the experiments studying the relationship between γ_{SO_2} and particle-phase
124 peroxide content, the molar ratio of organic peroxide to ammonium sulfate (Expt. 10-12, 15-18)
125 in the solution was adjusted to be 0.02, 0.2, 1, 2, and 4, respectively. In experiments where
126 malonic acid was used (Expt. 19-22), molar ratios of 0.2, 1, 2, and 4 were adopted. For
127 measuring γ_{SO_2} with different aerosol pH (Expt. 17, 23-25), different amounts of HCl (37%,
128 Sigma-Aldrich) were added into the solution prior to atomization. The atomized particles were
129 flowed into the chamber without drying, and therefore assumed to remain deliquesced under the
130 range of RH we studied. Expt. 2-14 also represent those where the relationship between γ_{SO_2} and
131 RH conditions were studied.

132 In the second set of experiments, the uptake of SO_2 onto SOA was investigated (Fig. S2, Expt.
133 26-28). A custom-built 10 L quartz oxidation flow reactor was used to produce SOA (Ye et al.,
134 2016) from different hydrocarbon precursors. In this work, we studied SOA formed from toluene
135 photooxidation, limonene ozonolysis and α -pinene ozonolysis, 3 of the most commonly studied
136 SOA systems (Ng et al., 2007; Hildebrandt et al., 2009; Hartz et al., 2005; Varutbangkul et al.,
137 2006). Toluene (analytical standard, Sigma Aldrich) was injected continuously into zero air flow



138 by a syringe (1000 mL, Hamilton) installed on a syringe pump (KDS Legato100) to achieve an
139 initial concentration of 0.5 ppm. Limonene (Sigma-Aldrich, 97 %) and α -pinene (Sigma-Aldrich,
140 98 %) were pre-dissolved in cyclohexane (Sigma-Aldrich, 99.5 %) with a volumetric ratio of 1:
141 1500 and 1: 500 to ensure that OH formed from limonene or α -pinene ozonolysis is scavenged
142 by cyclohexane, estimated based on the rate constants (Atkinson and Arey, 2003). The initial
143 steady-state concentrations of limonene and α -pinene were controlled to be around 2 ppm and 1
144 ppm entering the flow tube. O₃, used as the oxidant (for limonene and α -pinene) or the OH
145 precursor (for toluene), was generated by passing 0.5 L min⁻¹ pure oxygen (99.6 %, Linde,
146 Mississauga, Canada) through an O₃ generator (no. 97006601, UVP, Cambridge, UK).
147 Humidified air was produced by bubbling zero air through a custom-made humidifier at a flow
148 rate of 1 L min⁻¹. The photolysis of O₃ produces O (¹D), which reacts with water vapour to
149 produce ·OH with illumination from the 254 nm UV lamps (UVP, Cambridge, UK) to initiate
150 the photooxidation of toluene. The average residence time inside the flow tube was controlled to
151 be around 5 minutes. A gas chromatography–flame ionization detector (GC-FID, model 8610C,
152 SRI Instruments Inc., LV, USA) equipped with a Tenax® TA trap was used to monitor the
153 concentration of hydrocarbon precursors at the inlet/outlet of the flow reactor. In all cases, the O₃
154 concentration was maintained to be at least 10 times higher than that of the hydrocarbon.
155 Temperature and relative humidity were monitored by an Omega HX94C RH/T transmitter.
156 Particle size distribution and volume concentration were monitored using a custom-built
157 scanning mobility particle sizer (SMPS), which is a combination of a differential mobility
158 analyzer column (DMA, model 3081, TSI, Shoreview, MN, USA) with flow controls and a
159 condensation particle counter (CPC, model 3772, TSI, Shoreview, MN, USA).
160



161 2.2 Quantification of γ_{SO_2}

162 Prior to each experiment, the chamber was flushed by purified air overnight with a flow rate of
163 25 L min^{-1} until particle number concentration was less than 5 cm^{-3} and SO_2 was less than 1 ppb.
164 To adjust RH, the chamber was humidified by passing purified air through a custom-built
165 humidifier filled with ultra-pure water. For experiments with atomized ammonium sulfate or
166 malonic acid, SO_2 was injected into the chamber prior to the introduction of particles. For
167 experiments studying γ_{SO_2} onto SOA, aerosol generated from the flow tube was injected into the
168 Teflon chamber continuously after passing through an O_3 denuder (Ozone Solutions, Iowa, USA)
169 to achieve specific aerosol concentration inside the chamber prior to SO_2 addition. SO_2 mixing
170 ratio in the chamber during each experiment was continuously monitored using an SO_2 analyzer
171 (Model 43i, Thermo Scientific). The initial mixing ratio of SO_2 in each experiment was
172 controlled to be around 200 ppb. Aerosol size distribution was monitored by SMPS. The reactive
173 uptake coefficient of SO_2 was calculated by integrating the following equation:

$$174 \quad -\frac{d[\text{SO}_2]}{dt} = \frac{1}{4} \gamma_{\text{SO}_2} A \bar{c} [\text{SO}_2] \quad (1)$$

175
176 Where $[\text{SO}_2]$ is the SO_2 mixing ratio (ppb) monitored by the SO_2 analyzer; A is the average
177 surface area concentration ($\mu\text{m}^2 \text{ cm}^{-3}$) derived from the particle size distribution measured by
178 SMPS; \bar{c} represents the mean molecular velocity (cm s^{-1}) of SO_2 . $d[\text{SO}_2]/dt$ is solved over the
179 initial SO_2 decay, such that the peroxide concentration in the aerosol liquid phase is assumed to
180 be constant and pseudo-first order kinetics can be applied (Abbatt et al., 2012; Thornton et al.,
181 2003). A summary of all the measured γ_{SO_2} can be found in Table S1. Typical evolution of
182 monitored species can be seen in Fig. 1. Control experiments were performed in order to rule out
183 other potential factors (e.g. SO_2 loss in the in-line filter in front of the SO_2 analyzer, interferences
184 inside the SO_2 analyzer, chamber wall losses, SO_2 uptake onto wet ammonium sulfate, gas-phase



185 reaction of SO₂ with peroxide vapour) that may contribute to the SO₂ decay observed during the
186 γ_{SO_2} measurement inside the chamber (Fig. S3-S6). Measurement uncertainty of γ_{SO_2} in this study
187 was estimated from Expt. 10-12 to be 26%. Also, we observed there was SO₂ repartitioning from
188 the humid chamber wall in the presence of organic peroxide under high RH (Fig. S6b, RH 74%).
189 The observed SO₂ repartitioning rate was then applied to correct the γ_{SO_2} measured under high
190 RH conditions (above 70%, Expt.14), and this correction amounts to a 40% increase in
191 calculated γ_{SO_2} .

192

193 **2.3 Offline peroxide quantification**

194 Aerosol was collected onto 47 mm PTFE (polytetrafluoroethylene) filters with 0.2 μm pore size
195 (Whatman®, GE Healthcare) from the chamber by a diaphragm pump (KNF Neuberger Inc.,
196 USA) for offline chemical analysis. The particulate peroxide content in these samples prior to
197 SO₂ uptake was quantified using the iodometric–spectrophotometric assay (Docherty et al.,
198 2005). I₂ produced from the reaction between I⁻ and peroxides can further quickly combine with
199 the excess amount of I⁻ to form I₃⁻, which has brown color and absorbs UV-vis at 470nm. The
200 SOA extraction was then aliquoted into a 96-well UV plate (Greiner Bio-One, Kremsmünster,
201 AT) with 160 μL well⁻¹. 20 μL of formic acid ($\geq 95\%$, Sigma-Aldrich) was added into each
202 well, following by 20 μL potassium iodide (BioUltra, $\geq 99.5\%$, Sigma-Aldrich) solution
203 (dissolved in DI water). The plate was then covered by an adhesive plate sealer (EdgeBio,
204 Gaithersburg, USA) immediately in order to avoid reagent evaporation and O₂ oxidation. After
205 incubation for an hour in the dark, the UV-vis absorption at 470nm was measured using a UV-
206 vis spectrophotometer (Spectramax 190, Molecular Devices Corporation, Sunnyvale, CA) and
207 then converted to peroxide concentration using the calibration curve made by tert-butyl



208 hydroperoxide (70 wt. % in H₂O, Sigma-Aldrich) with a series of concentrations (0-10mM). The
209 average molecular mass for aerosol was assumed based on the chemical composition in order to
210 calculate the molar fraction of total peroxides. More details about the iodometric-
211 spectrophotometric procedures were described in previous work (Wang et al., 2018).

212

213 **3 Results and discussion**

214 **3.1 SO₂ uptake and RH**

215 A positive relationship between γ_{SO_2} and RH (between 25 and 71%) was observed for all types of
216 organic peroxides studied (Fig. 2). The positive dependence of the reactive uptake coefficient of
217 water-soluble gaseous species on RH has also been observed in other studies (Thornton et al.,
218 2003;Griffiths et al., 2009;Zhao et al., 2017;Zhang et al., 2019). Recently, the uptake behavior of
219 SO₂ onto soot, mineral dust and SOA were also shown to positively depend on RH (Zhang et al.,
220 2019;Zhao et al., 2017;Yao et al., 2019).

221 It is also noteworthy that an exponential dependence of SO₂ reactive uptake coefficient on RH
222 was observed in our study. γ_{SO_2} increases with increased relative humidity, which could even be
223 more significant under high RH regime. This is consistent with previous laboratory studies that
224 measured the reactive uptake coefficient of SO₂ onto aerosol to be exponentially dependent on
225 RH (Zhang et al., 2019;Yao et al., 2019). Additionally, multiple field campaigns have observed
226 significant correlation between particulate sulfate formation and ambient RH (Song et al.,
227 2019;Sun et al., 2013;Huang et al., 2020). Sun et al. (2013) observed faster sulfate formation rate
228 under humid conditions, proposing a significant impact of aerosol liquid water on sulfate
229 production during wintertime in Beijing. Zheng et al. (2015) reported a notably higher SOR
230 (molar ratio of sulfate to the sum of sulfate and SO₂) during wet period (RH>50%), indicating



231 the importance of heterogeneous reactions to the secondary sulfur transformation with abundant
232 aerosol water content under humid conditions. In a recent study by Song et al. (2019), the rapid
233 sulfate formation rate observed under high RH conditions was found to be significantly higher
234 than atmospheric modeling results implemented with homogeneous SO₂ oxidation pathways,
235 which was later attributed to heterogeneous sulfate formation mechanisms. Multiple mechanisms
236 can potentially explain this observed γ_{SO_2} -RH dependence. An enhanced relative humidity would
237 result in a nonlinear increase of aerosol water content, which can lead to more SO₂ dissolved in
238 the aerosol aqueous phase (Seinfeld and Pandis, 2012). It should be noted that while the relative
239 humidity is varied systematically in these experiments, the relationship is more complex since
240 RH also affects other aerosol properties which can affect the uptake kinetics in turn. For
241 example, a higher aerosol liquid water content could dilute protons and thus lower the aerosol
242 acidity. In a study by Laskin et al. (2003), an enhanced uptake of SO₂ onto sea-salt particles was
243 observed with an increased aerosol alkalinity at high pH range.

244

245 **3.2 Dependence of SO₂ uptake on peroxide content and type**

246 As expected, the measured uptake rate of SO₂ is dependent on the particulate peroxide content in
247 the current study. Fig. 3 shows that γ_{SO_2} is linearly proportional to the amount of particulate
248 peroxide for aerosol with similar volume-to-surface ratios and containing the same type of
249 organic peroxides. This positive relationship between γ_{SO_2} and condensed phase peroxide content
250 has also been inferred from experiments of SO₂ uptake onto α -pinene SOA (Yao et al., 2019),
251 where the peroxide content in α -pinene SOA was varied indirectly by introducing NO and
252 adjusting the branching ratio of the peroxide-yielding RO₂+HO₂/RO₂ pathway.



253 In addition to the amount of peroxide injected, the particulate fraction of organic peroxide
254 available for heterogeneous reaction is also influenced by gas-particle partitioning. As indicated
255 in Fig. 2, the reactive uptake coefficients of different organic peroxides vary amongst each other
256 by about an order of magnitude in the range of RH studied, despite the same amounts of peroxide
257 relative to ammonium sulfate initially in the atomizing solution. Based on our previous work
258 (Wang et al., 2019), the aqueous-phase rate constants for these organic peroxides with dissolved
259 S(IV) only vary by a factor of 2-3 and therefore cannot fully explain the observed difference in
260 uptake rates. Since vapour pressure vary considerably among the different peroxides in the
261 present study, gas-particle partitioning is likely to influence the amount of peroxide in the
262 particle phase that react with dissolved SO₂. The relative particulate peroxide content on filters
263 of the three peroxides collected from chamber experiments under RH 50% without SO₂ uptake
264 were measured by the offline KI method (Fig. S7). Although the initial ratio of organic peroxide
265 to ammonium sulfate in the atomizing solution was nominally the same, we measured the highest
266 amount of particulate peroxide with 2-butanone peroxide (16.7%), followed by cumene
267 hydroperoxide (12.7%) and then tert-butyl hydroperoxide (3.8%) using the offline iodometric
268 method. This trend in particulate peroxide content is consistent with the vapour pressures
269 calculated using the SIMPOL group contribution method (Pankow et al., 2008), with 2-butanone
270 peroxide being the least volatile, and tert-butyl hydroperoxide being the most volatile. Also, the
271 order of particle-phase peroxide content is consistent with the order of γ_{SO_2} observed, as shown in
272 Fig. 2. A simple visualization of these relationships between different peroxide characteristics
273 (number of peroxide groups, vapour pressure and aqueous-phase rate constants) and measured
274 γ_{SO_2} (at RH = 50%) is illustrated in Fig. S7, which indicates higher γ_{SO_2} can be expected for
275 multifunctional organic peroxides with lower vapour pressures and higher aqueous phase



276 reactivities. It should be noted that the order of magnitude difference in experimentally measured
277 γ_{SO_2} among various organic peroxides (Fig.2) is still not fully explained when both volatility and
278 reaction kinetics are taken into account (Fig.S7), suggesting that the reactive uptake may be
279 influenced by other factors. In summary, for our current experiments where we nominally
280 maintained total injected amount of organic peroxide constant, measured γ_{SO_2} depends both on
281 reactivity and gas-particle partitioning of the organic peroxides.

282

283 **3.3 SO₂ uptake and aqueous phase kinetics**

284 Since the aqueous phase reaction rate constants between S(IV) and these model organic
285 peroxides have been measured previously (Wang et al., 2019), we can test our understanding of
286 the measured γ_{SO_2} using a simple model. By assuming the amount of SO₂ dissolved in the aerosol
287 is in equilibrium with the gas phase, the overall γ_{SO_2} can be expressed using the simplified
288 resistor model (Hanson et al., 1994):

289

$$290 \quad \frac{1}{\gamma} = \frac{1}{\alpha} + \frac{\bar{c}}{4HRT\sqrt{k^1D_l}} \frac{1}{\left[\coth(q) - \frac{1}{q}\right]} \quad (2)$$

291 where α is the mass accommodation coefficient, \bar{c} is the mean molecular speed of SO₂ (cm s⁻¹),

292 H is the effective Henry's law constant that includes both the dissolution of SO₂ and the

293 dissociation of H₂SO₃ (M atm⁻¹), R is the ideal gas constant (atm L mol⁻¹ K⁻¹), T is the

294 temperature (K), and the parameter q is used to describe the competition between the reaction

295 and diffusion of the dissolved gaseous species within a particle, which is further calculated as:

$$296 \quad q = r \sqrt{\frac{k^1}{D_l}} \quad (3)$$

297 where r is the radius (cm) of a given particle, D_l is the aqueous-phase diffusion coefficient (cm²

298 s⁻¹), k¹ is the first order rate constant (s⁻¹) for the reaction. For experiments in the current study,



299 the calculated q values were consistently found to be far less than 1, which indicates a volume-
300 limited reaction regime. Combining with the assumption of a relatively fast mass
301 accommodation process compared with the bulk phase reaction, equation (2) can be further
302 simplified as to describes reactive uptake in the volume-limited regime:

$$303 \quad \gamma = \frac{4HRT[\text{peroxide}]k^{\text{II}}V}{\bar{c}S} \quad (4)$$

304 Here, we assume all the peroxides remain in the condensed phase upon atomizing and reaction
305 inside the chamber for the upper-bound prediction of γ_{SO_2} . [peroxide] represents the particle phase
306 concentration of total organic peroxide (M) based on the initial ratio between organic peroxide and
307 ammonium sulfate in the atomizing solution, and the aerosol water content output by E-AIM III
308 (Clegg et al., 1998), k^{II} is the second order reaction rate constant ($\text{M}^{-1} \text{s}^{-1}$), which we have
309 measured in the bulk phase at dilute concentrations previously (Wang et al., 2019), V/S is the ratio
310 between particle volume concentration ($\mu\text{m}^3 \text{cm}^{-3}$) and particle surface area concentration (μm^2
311 cm^{-3}) derived from SMPS measurements. As a result, the observed reactive uptake coefficient of
312 SO_2 can be compared to that predicted from the bulk phase reaction rate constant, and the results
313 are shown in Fig. 4 and Fig. S8. Overall, we noticed that this model captures the dependence of
314 γ_{SO_2} on peroxide content, but the modeled results were found to be generally 15-50 times lower
315 than the experimentally measured values (Fig. S8). The current γ_{SO_2} predictions are likely upper-
316 bound estimates since all the peroxides were assumed to stay in the condensed phase without
317 partitioning. As a result, this observed 15-50 times of discrepancy could even be larger if the
318 particulate peroxide content during the chamber experiments were lower due to partitioning.
319 It should be noted that the calculated γ_{SO_2} was based on reaction kinetics measured in dilute
320 solutions while the experimental γ_{SO_2} were measured directly from suspended particles. This large
321 difference in kinetics between those in aerosol and in dilute bulk solution suggests that this

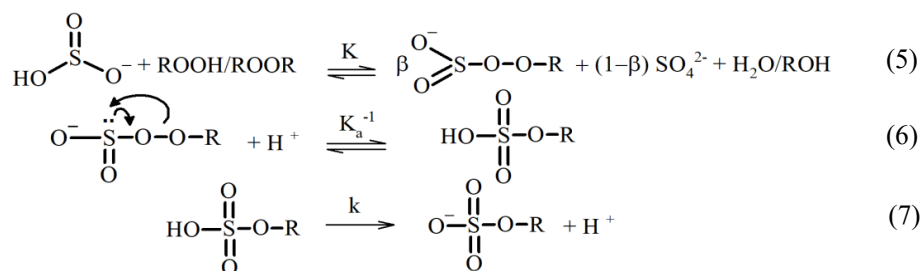


322 multiphase interaction is strongly favored in the highly concentrated aerosol environment. One of
323 the potential explanations for this discrepancy could be liquid-liquid phase separation (LLPS) in
324 aerosol between organic peroxide and ammonium sulfate (Ciobanu et al., 2009; O'Brien et al.,
325 2015) such that SO₂ can directly interact with the acidic organic phase, where the concentration of
326 peroxides can be higher and the kinetics can be different from what we have measured in dilute
327 solution (Wang et al., 2019). However, LLPS is generally governed by the chemical composition
328 of the hydrophobic phase (Freedman, 2017). A higher level of oxygenation in organic aerosol is
329 related with higher hydrophilicity, which would favor a homogeneous particle instead of phase
330 separation. Previous studies showed that LLPS did not occur for organic coating with O:C above
331 0.8 (You et al., 2013; You et al., 2014). The LLPS phenomenon in simple organic-inorganic
332 mixtures can also be affected by the functional groups. The maximum O:C for LLPS could be 0.71
333 for organics with multiple carboxylic and hydroxyl groups but low aromatic content (Song et al.,
334 2012) while the 2-butanone peroxide we used for both γ_{SO_2} measurement and prediction in the
335 present study has multiple peroxide groups with an O:C value of 0.75. Particle size could also have
336 impacts on phase separation (Cheng et al., 2015). Particle diameters in the current study are mainly
337 under 200 nm while a previous study showed particles smaller than this size are less likely to
338 experience LLPS (Veghte et al., 2013). We therefore believe that LLPS is not likely to be
339 responsible for the enhanced uptake rate observed under these experimental conditions.

340 Another explanation is the high solute strength in the concentrated aerosol phase. As indicated in
341 Fig. 4 and Fig. S8, the difference between the measured and predicted γ_{SO_2} is larger for ammonium
342 sulfate aerosol than for malonic acid. Meanwhile, the calculated ionic strength in aerosol liquid
343 phase under RH 50% for ammonium sulfate (40 mol kg⁻¹) is significantly larger than that of
344 malonic acid (0.45 mol kg⁻¹). It has been previously reported that the reaction rate between sulfite



345 and hydrogen peroxide in aqueous phase increases with ionic strength (Maaß et al., 1999). Based
346 on the reaction mechanisms proposed for dissolved SO_2 and hydrogen peroxide (Halperin and
347 Taube, 1952), we speculate the reaction between aqueous phase S(IV) and organic peroxides to
348 follow a similar mechanism:



349 where the overall rate constant is equal to $k \frac{K}{K_a}$, assuming fast equilibrium steps for reactions 5 and
350 6. Dissociated solutes are surrounded by an extended solvation shell which could affect the
351 reaction rates (Herrmann, 2003). Fewer available free water molecules would therefore shift the
352 equilibrium to the right in equation (5). Additionally, higher ionic strength also corresponds to an
353 increased concentration of electrolytes in the aqueous phase, which could hinder the dissociation
354 of the peroxymonosulfurous acid and shift the equilibrium in equation (6) to the right. In recent
355 work by Liu et al. (2020), the rate of S(IV) oxidation by H_2O_2 can be enhanced by up to a factor
356 of 50 in aerosol aqueous phase compared to that of dilute solution. The highest ionic strength at
357 which such enhancement was measured for the H_2O_2 oxidation pathway was 15 mol kg^{-1} (Liu et
358 al., 2020).

359 Whereas the above analysis is based on the assumption that all the chemistry occurs in the bulk
360 component of the particle, it is also possible that some component of the reaction occurs at the gas-
361 particle interface and the overall kinetics can be affected by interfacial characteristics. For example,
362 an enhanced ionic strength in the aerosol phase can also impact the interfacial reaction mechanisms.
363 Previous study has shown evidence that interfacial chemistry is important for SO_2 oxidation in the



364 aerosol phase (Laskin et al., 2003). With higher ionic strength, anions partitioning to the air-liquid
365 interface can promote the overall reaction kinetics via proton transfer and thus accelerate the
366 interfacial chemistry (Knipping et al., 2000; Mishra et al., 2012; Mekic et al., 2018; Mekic et al.,
367 2020; Wei et al., 2018; Ruiz-Lopez et al., 2020). However, it should be noted that there is no
368 evidence from the current study showing direct relationship between the interfacial properties and
369 γ_{SO_2} , and future studies are warranted.

370 Therefore, while more studies are needed to clearly delineate the roles of ionic strength, interfacial
371 activity, bulk reactivity, and particle phase state quantitatively, the enhancement of SO_2 oxidation
372 kinetics by highly concentrated aerosol particles compared to dilute aqueous solutions are
373 concluded to be large (factor of 15-50) for the experimental conditions in the current study.

374

375 **3.4 SO_2 uptake and aerosol pH**

376 As indicated by the proposed reaction mechanisms (Eqn. 5-7), protons are important reaction
377 intermediates for this SO_2 oxidation pathway. Previously, the aqueous phase reaction rate
378 constants between organic peroxides and dissolved SO_2 were measured to be pH dependent
379 (Wang et al., 2019). Moreover, the dissolution equilibrium of SO_2 into aqueous phase is also pH
380 sensitive (Seinfeld and Pandis, 2012). Besides, many studies have shown that the uptake kinetics
381 for gaseous species can be affected by the condensed phase pH (Shi et al., 1999; Gaston et al.,
382 2014; Drozd et al., 2013; Jang and Kamens, 2001; Liu et al., 2015). Reactive uptake of ammonia
383 was observed to depend on condensed phase acidity (Shi et al., 1999). Heterogeneous
384 condensation of isoprene-derived epoxydiol onto seed aerosol was found to increase with proton
385 concentration (Gaston et al., 2014). In the current study, the potential impact from particle phase
386 pH on γ_{SO_2} was explored by adding HCl into the atomizing solution. To estimate the particle



387 phase pH, two different methods associated with two different assumptions were used. In the
388 first scenario, the aerosol pH in each experiment was estimated using the E-AIM III model
389 (Clegg et al., 1998) based on the initial molar ratios of inorganic species (H^+ , NH_4^+ , SO_4^{2-} , Cl^-) in
390 the atomizing solution and measured RH (around 50%). In the second scenario, the additional
391 sulfate formed from reactive uptake of SO_2 was taken into consideration. The partitioning of HCl
392 was allowed in the model simulation for both scenarios. The formation of sulfate would enhance
393 the proton concentration in the aerosol liquid phase thus lower the aerosol pH. The average pH
394 during the SO_2 uptake process is likely in between these two extremes.

395 Fig. 5 shows the measured reactive uptake coefficients of SO_2 as a function of the calculated pH.
396 The reactive uptake coefficient was found to increase with increasing proton concentrations
397 (decreasing pH), which is consistent with acid-catalyzed reactions between peroxides and
398 dissolved SO_2 as measured in the bulk phase (Lind et al., 1987; Wang et al., 2019). γ_{SO_2} was also
399 predicted for the same range of pH based on Eqn. 4 and the pH-dependent bulk-phase reaction
400 rate constants measured previously (Wang et al., 2019). Indicated by Fig. 5, the measured γ_{SO_2}
401 again exceeds the predicted γ_{SO_2} by about a factor of 50, which is consistent with what we
402 reported earlier and is likely due to the effects of aerosol ionic strength.

403 Unlike the observed γ_{SO_2} , however, the predicted γ_{SO_2} does not exhibit a monotonic trend. γ_{SO_2} is
404 expected to decrease with decreasing pH at high pH (>2) as the effective Henry's law constant of
405 SO_2 decreases with higher acidity (Seinfeld and Pandis, 2012). γ_{SO_2} is not expected to increase
406 with decreasing pH until pH is below 2 where the acidity enhancement in reaction rate constant
407 exceeds the decrease in SO_2 solubility. As illustrated earlier, extrapolating dilute aqueous-phase
408 kinetics to the highly concentrated aerosol requires considering effects from high solute strength.
409 Solute strength may change the pH dependence of γ_{SO_2} in two ways. First, the solubility of SO_2



410 may decrease and become less dependent on pH as ionic strength increases (Rodríguez-Sevilla et
411 al., 2002). A former study (Leng et al., 2015) has shown that the effective Henry's law of
412 triethylamine decreases with increased ionic strength. Another potential explanation is that the
413 aqueous phase reaction rate constant can be more pH-dependent at high ionic strengths than what
414 we measured previously in dilute solutions. In either case, the inflection of the predicted γ_{SO_2}
415 would change and γ_{SO_2} could become more negatively dependent on pH ($d[\gamma_{\text{SO}_2}]/d[\text{pH}]$ becomes
416 less positive in the high pH range and/or more negative in the low pH range), which would
417 match more closely with the observed dependence. It should also be noted that there are
418 substantial uncertainties in estimating pH values since the reactive uptake is a dynamic process
419 and will influence aerosol pH in turn upon sulfate formation. In summary, while the magnitude
420 of predicted γ_{SO_2} is consistent with our expected values (after accounting for the enhancement by
421 high aerosol solute strength), we cannot fully explain the dependence of γ_{SO_2} on aerosol pH at the
422 current stage. Future studies should investigate how the effective Henry's law of SO_2 and pH
423 dependence of reaction rate constants vary in aerosol liquid phase with high solute strength in
424 order to have a more comprehensive understanding of the relationship between γ_{SO_2} and aerosol
425 pH.

426 **3.4 SO_2 uptake onto SOA**

427 γ_{SO_2} was measured for a few model SOA systems, as organic peroxides are abundant in SOA
428 (Surratt et al., 2006;Kautzman et al., 2009;Krapf et al., 2016;Bonn et al., 2004). Here we studied
429 SOA formed from monoterpene ozonolysis and toluene photooxidation. It should be noted that
430 for the γ_{SO_2} measurements of toluene SOA, a strong hydrocarbon interference was observed with
431 the SO_2 analyzer, likely stemming from the high concentrations of gas-phase aromatic
432 compounds. A rough estimate of the uptake rate for toluene SOA from aerosol mass



433 spectrometer sulfate measurements is provided in the SI (Section 1). The reactive uptake
434 coefficient of SO₂ onto Saharan mineral dust was reported on the order of 10⁻⁵ (Adams et al.,
435 2005). γ_{SO_2} onto dust with the coexistence of NO₂ and NH₃ under various RH conditions were
436 measured to be 10⁻⁷ to 10⁻⁵ (Zhang et al., 2019). For a variety of metal oxides, SO₂ reactive
437 uptake coefficients were quantified to be between 10⁻⁶ and 10⁻⁴ (Usher et al., 2002; Fu et al.,
438 2007; Shang et al., 2010). More recently, γ_{SO_2} studied for heterogeneous sulfate formation by
439 photolysis of particulate nitrate were reported in the range of 10⁻⁶ to 10⁻⁵ (Gen et al., 2019). As
440 shown in Fig. 6, γ_{SO_2} for all SOA systems were measured to be on the order of 10⁻⁴. Similar γ_{SO_2}
441 values on the order of 10⁻⁴ were measured for α -pinene SOA by Yao et al. (2019), and 10⁻⁵ for
442 limonene SOA estimated from the chamber study by Ye et al. (2018). The reaction products
443 from this SOA and SO₂ interaction will be reported in a separate study.

444

445 **4. Atmospheric Implications**

446 Oxidation of atmospheric hydrocarbons produces reactive intermediates that can potentially
447 interact with SO₂ and form particulate sulfate, contributing to PM formation and growth (Berndt
448 et al., 2015; Mauldin et al., 2012; Yao et al., 2019). Organic peroxides generated from both
449 biogenic and anthropogenic hydrocarbon emissions are abundant in submicron aerosol. Given
450 that they are highly reactive with relatively short lifetimes (Bonn et al., 2004; Krapf et al.,
451 2016; Qiu et al., 2020), these species could serve as important condensed phase oxidants for gas
452 phase SO₂. Combining laboratory measurements and model predictions, the current study
453 investigated heterogeneous reactions between SO₂ and particulate organic peroxide. The
454 measured γ_{SO_2} for organic peroxide containing aerosol ranges from 10⁻⁵ to 10⁻² in this study.
455 Based on the modeling work by Wang et al. (2014), adding an SO₂ uptake pathway to GEOS-



456 Chem with a reactive uptake coefficient of 10^{-4} could improve the surface sulfate prediction by
457 more than 50% during a severe haze episode over North China (RH 50%), suggesting the
458 potential importance of this multiphase reaction pathway, especially when SOA is the dominant
459 component in particulate matter.

460 The dependence of the heterogeneous kinetics on RH, aerosol pH, peroxide type, and peroxide
461 content were also evaluated. The experimentally measured γ_{SO_2} was found to be consistently
462 higher than that predicted from reaction kinetics with organic peroxides in the dilute aqueous
463 phase. This discrepancy can be potentially explained by the effects of high ionic strength
464 presented in the aerosol, suggesting that the impact from highly concentrated solutes needs to be
465 taken into consideration when applying aqueous phase kinetics to aerosol multiphase chemistry,
466 especially for particles containing strong electrolytes. We also observed that the kinetics of this
467 multiphase reaction exhibit a weak dependence on pH. Increasing the condensed-phase acidity
468 enhances the heterogeneous rate constant at low pH range, and while this pH dependence is
469 consistent with that of the aqueous phase reaction rate constant measured previously, it is not
470 consistent with the decrease of effective Henry's law constant of SO_2 along with enhanced
471 acidity. Currently, we are not able to fully explain the pH dependence, likely due to the
472 uncertainties from high ionic strength, and further studies are warranted. Particle phase peroxide
473 content was observed to be linearly correlated with γ_{SO_2} . Moreover, γ_{SO_2} measured for 2-butanone
474 peroxide was found to be orders of magnitude higher than that of cumene hydroperoxide and
475 tert-butyl hydroperoxide. The difference in γ_{SO_2} among various types of organic peroxides can be
476 partially explained by their condensed-phase reactivity and gas-particle partitioning.

477 In general, we found the observed γ_{SO_2} in this study can be summarized using the following
478 semiempirical multilinear relationship:



479
$$\log \gamma = -1.7 + 0.0024 \times k'' + 0.46 \times PAS + 0.024 \times RH - 1.9 \times Vp \quad (8)$$

480 where γ is the reactive uptake coefficient, k'' is the aqueous phase S(IV) oxidation rate constant
481 ($\text{M}^{-1} \text{s}^{-1}$), PAS is the molar ratio between particulate peroxide and ammonium sulfate in the
482 atomizing solution, which is a proxy for the amount of peroxide in both gas and particle phases
483 applied in the current study, RH is the relative humidity (%), Vp is the vapour pressure (kPa) of
484 the peroxide. Fig. 7 illustrates the degree to which this semi-empirical expression describes the
485 experimental data for ammonium sulfate aerosol mixed with the three types of organic peroxides.
486 Residual evaluations of this multilinear regression can be found Fig. S9. We caution that this
487 equation is not directly applicable to atmospheric models in its current form, especially since the
488 particle phase peroxide content (PAS) value we applied as input is a calculated value, rather than
489 a measurement. However, it illustrates the internal consistency of our experimental results across
490 a range of RH, peroxide content, and aqueous phase reactivities, which are the key variables for
491 uptake rates. Better understanding of ionic strengths and pH in aerosol, either through modeling
492 or direct measurements of these variables, is needed to establish the coefficient dependence.
493 Future studies should be focused on exploring γ_{SO_2} and the reaction products for various types of
494 SOA as well as ambient particles under atmospherically relevant conditions, evaluating the
495 underlying impacts from photochemical condition, chemical composition, particle morphology,
496 ionic strength and interfacial properties on this multiphase physicochemical process. Overall,
497 γ_{SO_2} presented in our study and its relationship with ambient RH, aerosol pH, ionic strength,
498 particulate peroxide content and type could provide a framework for the implementation of this
499 heterogeneous mechanism in atmospheric models to have a better understanding of ambient
500 sulfate formation and particle growth.

501



502 *Author contributions*

503 A.W.H. C. and S.W. designed the study. S.W., T. L., and J. J. performed the experiments. S.W.,
504 A.W.H. C., T. L., and J. J. analyzed data. S.W. and A.W.H. C. wrote the manuscript with the
505 input from all co-authors.

506

507 *Data availability*

508 All data presented in this study are available in the supplemental material and have been
509 deposited in figshare.

510

511 *Associated content*

512 Supporting Information.

513

514 *Competing interests*

515 The authors declare no competing financial interest.

516

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526 Reference

- 527 Abbatt, J. P. D., Lee, A. K. Y., and Thornton, J. A.: Quantifying trace gas uptake to tropospheric
528 aerosol: recent advances and remaining challenges, *Chem. Soc. Rev.*, 41, 6555–6581,
529 <https://doi.org/10.1039/c2cs35052a>, 2012.
- 530 Adams, J. W., Rodriguez, D., and Cox, R. A.: The uptake of SO₂ on Saharan dust: a flow tube
531 study, *Atmos. Chem. Phys.*, 5, 2643–2676, doi:10.5194/acp-5-2643-2005, 2005.
- 532 Andreae, M.O., Afchine, A., Albrecht, R., Holanda, B.A., Artaxo, P., Barbosa, H.M., Borrmann,
533 S., Cecchini, M.A., Costa, A., Dollner, M. and Fütterer, D.: Aerosol characteristics and particle
534 production in the upper troposphere over the Amazon Basin, *Atmos. Chem. Phys.*, 18, 921–961,
535 2018.
- 536 Atkinson, R., and Arey, J.: Atmospheric degradation of volatile organic compounds, *Chem.*
537 *Rev.*, 103, 4605–4638, 2003.
- 538 Berndt, T., Richters, S., Kaethner, R., Voigtländer, J., Stratmann, F., Sipilä, M., Kulmala, M.,
539 and Herrmann, H.: Gas-phase ozonolysis of cycloalkenes: Formation of highly oxidized RO₂
540 radicals and their reactions with NO, NO₂, SO₂, and other RO₂ radicals, *J. Phys. Chem. A*, 119,
541 10336–10348, 10.1021/acs.jpca.5b07295, 2015.
- 542 Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crouse, J.
543 D., Wennberg, P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop,
544 D. R., Thornton, J. A., Donahue, N., Kjaergaard, H. G., and Ehn, M.: Highly oxygenated organic
545 molecules (HOM) from gas-phase autoxidation involving peroxy radicals: A key contributor to
546 atmospheric aerosol, *Chem. Rev.*, 119, 3472–3509, 10.1021/acs.chemrev.8
547 b00395, 2019.
- 548 Bonn, B., von Kuhlmann, R., and Lawrence, M. G.: High contribution of biogenic
549 hydroperoxides to secondary organic aerosol formation, *Geophys. Res. Lett.*, 31, L10108,
550 <https://doi.org/10.1029/2003GL019172>, 2004.
- 551 Chen, Q., Farmer, D. K., Schneider, J., Zorn, S. R., Heald, C. L., Karl, T. G., Guenther, A.,
552 Allan, J. D., Robinson, N., Coe, H., Kimmel, J. R., Pauliquevis, T., Borrmann, S., Pöschl, U.,
553 Andreae, M. O., Artaxo, P., Jimenez, J. L., and Martin, S. T.: Mass spectral characterization of
554 submicron biogenic organic particles in the Amazon Basin, *Geophys. Res. Lett.*, 36, L20806,
555 <https://doi.org/10.1029/2009GL039880>, 2009.
- 556 Chen, T., Chu, B., Ge, Y., Zhang, S., Ma, Q., He, H., and Li, S.-M.: Enhancement of aqueous
557 sulfate formation by the coexistence of NO₂/NH₃ under high ionic strengths in aerosol water,
558 *Environ. Pollut.*, 252, 236–244, <https://doi.org/10.1016/j.envpol.2019.05.119>, 2019.
- 559 Cheng, Y., Su, H., Koop, T., Mikhailov, E., and Pöschl, U.: Size dependence of phase transitions
560 in aerosol nanoparticles, *Nat. Commun.*, 6, 5923, 10.1038/ncomms6923, 2015.
- 561 Cheng, Y. F., Zheng, G. J., Wei, C., Mu, Q., Zheng, B., Wang, Z. B., Gao, M., Zhang, Q., He, K.
562 B., Carmichael, G., Pöschl, U., and Su, H.: Reactive nitrogen chemistry in aerosol water as a
563 source of sulfate during haze events in China, *Sci. Adv.*, 2, e1601530, <https://doi.org/10.1126/sciadv.1601530>, 2016.
- 564 Ciobanu, V. G., Marcolli, C., Krieger, U. K., Weers, U., and Peter, T.: Liquid-liquid phase
565 separation in mixed organic/inorganic aerosol particles, *J. Phys. Chem. A*, 113, 10966–10978,
566 2009.
- 567 Clegg, S. L., Brimblecombe, P., and Wexler, A. S.: Thermodynamic model of the system
568 H⁺–NH₄⁺–Na⁺–SO₄²⁻–NO₃⁻–Cl⁻–H₂O at 298.15 K, *J. Phys. Chem. A*, 102, 2155–2171,
569 <https://doi.org/10.1021/jp973043j>, 1998.
- 570



- 571 Docherty, K. S., Wu, W., Lim, Y. B., and Ziemann, P. J.: Contributions of organic peroxides to
572 secondary aerosol formed from reactions of monoterpenes with O₃, *Environ. Sci. Technol.*, 39,
573 4049-4059, 2005.
- 574 Dovrou, E., Rivera-Rios, J. C., Bates, K. H., and Keutsch, F. N.: Sulfate Formation via Cloud
575 Processing from Isoprene Hydroxyl Hydroperoxides (ISOPOOH), *Environ. Sci. Technol.*, 53,
576 12476-12484, [10.1021/acs.est.9b04645](https://doi.org/10.1021/acs.est.9b04645), 2019.
- 577 Drozd, G. T., Woo, J. L., and McNeill, V. F.: Self-limited uptake of α -pinene oxide to acidic
578 aerosol: the effects of liquid-liquid phase separation and implications for the formation of
579 secondary organic aerosol and organosulfates from epoxides, *Atmos. Chem. Phys.*, 13, 8255–
580 8263, [doi:10.5194/acp-13-8255-2013](https://doi.org/10.5194/acp-13-8255-2013), 2013.
- 581 Epstein, S. A., Blair, S. L., and Nizkorodov, S. A.: Direct photolysis of α -pinene ozonolysis
582 secondary organic aerosol: effect on particle mass and peroxide content, *Environ. Sci. Technol.*,
583 48, 11251-11258, 2014.
- 584 Fairlie, T. D., Jacob, D. J., Dibb, J. E., Alexander, B., Avery, M. A., van Donkelaar, A., and
585 Zhang, L.: Impact of mineral dust on nitrate, sulfate, and ozone in transpacific Asian pollution
586 plumes, *Atmos. Chem. Phys.*, 10, 3999–4012, [doi:10.5194/acp-10-3999-2010](https://doi.org/10.5194/acp-10-3999-2010), 2010.
- 587 Freedman, M. A.: Phase separation in organic aerosol, *Chem. Soc. Rev.*, 46, 7694–7705,
588 <https://doi.org/10.1039/C6CS00783J>, 2017.
- 589 Fu, H. B., Wang, X., Wu, H. B., Yin, Y., and Chen, J. M.: Heterogeneous uptake and oxidation
590 of SO₂ on iron oxides, *J. Phys. Chem. C*, 111, 6077–6085, 2007.
- 591 Gao, M., Carmichael, G. R., Wang, Y., Saide, P. E., Yu, M., Xin, J., Liu, Z., and Wang, Z.:
592 Modeling study of the 2010 regional haze event in the North China Plain, *Atmos. Chem. Phys.*,
593 16, 1673–1691, <https://doi.org/10.5194/acp-16-1673-2016>, 2016.
- 594 Gaston, C. J., Riedel, T. P., Zhang, Z., Gold, A., Surratt, J. D., and Thornton, J. A.: Reactive
595 uptake of an isoprene-derived epoxydiol to submicron aerosol particles, *Environ. Sci. Technol.*,
596 48, 11178-11186, [10.1021/es5034266](https://doi.org/10.1021/es5034266), 2014.
- 597 Gen, M., Zhang, R., Huang, D. D., Li, Y., and Chan, C. K.: Heterogeneous SO₂ Oxidation in
598 Sulfate Formation by Photolysis of Particulate Nitrate, *Environ. Sci. Tech. Lett.*, 6, 86–91,
599 <https://doi.org/10.1021/acs.estlett.8b00681>, 2019.
- 600 Griffiths, P. T., Badger, C. L., Cox, R. A., Folkers, M., Henk, H. H., and Mentel, T. F.: Reactive
601 uptake of N₂O₅ by aerosols containing dicarboxylic acids. Effect of particle phase, composition,
602 and nitrate content, *J. Phys. Chem. A*, 113, 5082-5090, [10.1021/jp8096814](https://doi.org/10.1021/jp8096814), 2009.
- 603 Gunz, D. W. and Hoffmann, M. R.: Atmospheric chemistry of peroxides: A review, *Atmos.*
604 *Environ.*, 24A, 1601–1633, [https://doi.org/10.1016/0960-1686\(90\)90496-A](https://doi.org/10.1016/0960-1686(90)90496-A), 1990.
- 605 Guo, H., Weber, R. J., and Nenes, A.: High levels of ammonia do not raise fine particle pH
606 sufficiently to yield nitrogen oxide-dominated sulfate production, *Sci. Rep.*, 7, 12109, 2017.
- 607 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen,
608 J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T.,
609 Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W.,
610 McFiggans, G., Mentel, Th. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D.,
611 Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic
612 aerosol: current and emerging issues, *Atmos. Chem. Phys.*, 9, 5155–5236, <https://doi.org/10.5194/acp9-5155-2009>, 2009.
- 614 Halperin, J., and Taube, H.: The transfer of oxygen atoms in oxidation—reduction reactions. IV.
615 The reaction of hydrogen peroxide with sulfite and thiosulfate, and of oxygen, manganese
616 dioxide and of permanganate with sulfite, *J. Am. Chem. Soc.*, 74, 380-382, 1952.



- 617 Hanson, D. R., Ravishankara, A. R., and Solomon, S.: Heterogeneous reactions in sulfuric acid
618 aerosols: A framework for model calculations, *J. Geophys. Res.*, 99, 3615, [https://doi.org/](https://doi.org/10.1029/93JD02932)
619 10.1029/93JD02932, 1994.
- 620 Hartz, K. E. H., Rosenorn, T., Ferchak, S. R., Raymond, T. M., Bilde, M., Donahue, N. M., and
621 Pandis, S. N.: Cloud condensation nuclei activation of monoterpene and sesquiterpene
622 secondary organic aerosol, *J. Geophys. Res.-Atmos.*, 110(D14), D14208, doi:10.1029/2004J
623 D005754, 2005.
- 624 He, L.-Y., Huang, X.-F., Xue, L., Hu, M., Lin, Y., Zheng, J., Zhang, R., and Zhang, Y.-H.:
625 Submicron aerosol analysis and organic source apportionment in an urban atmosphere
626 in Pearl River Delta of China using high-resolution aerosol mass spectrometry, *J. Geophys. Res.-*
627 *Atmos.*, 116, D12304, <https://doi.org/10.1029/2010JD014566>, 2011.
- 628 Herrmann, H.: Kinetics of aqueous phase reactions relevant for atmospheric chemistry, *Chem.*
629 *Rev.*, 103, 4691–4716, 2003.
- 630 Hildebrandt, L., Donahue, N. M., Pandis, S. N.: High formation of secondary organic aerosol
631 from the photo-oxidation of toluene, *Atmos. Chem. Phys.*, 9, 2973–2986, [https://doi.org](https://doi.org/10.5194/acp-9-2973-2009)
632 /10.5194/acp-9-2973-2009, 2009.
- 633 Huang, L., An, J., Koo, B., Yarwood, G., Yan, R., Wang, Y., Huang, C., and Li, L.: Sulfate
634 formation during heavy winter haze events and the potential contribution from heterogeneous
635 SO₂ + NO₂ reactions in the Yangtze River Delta region, China, *Atmos. Chem. Phys.*, 19, 14311–
636 14328, 10.5194/acp-19-14311-2019, 2019.
- 637 Huang, L., Coddens, E. M., and Grassian, V. H.: Formation of organosulfur compounds from
638 aqueous phase reactions of S (IV) with methacrolein and methyl vinyl ketone in the presence of
639 transition metal ions, *ACS Earth Space Chem.*, 3, 1749–1755, 2019.
- 640 Huang, L., Liu, T. and Grassian, V. H.: Radical-initiated formation of aromatic organosulfates
641 and sulfonates in the aqueous phase. *Environ. Sci. Technol.*, 54, 11857–11864,
642 <https://doi.org/10.1021/acs.est.0c05644>, 2020.
- 643 Huang, R. J., Zhang, Y. L., Bozzetti, C., Ho, K. F., Cao, J. J., Han, Y. M., Daellenbach, K. R.,
644 Slowik, J. G., Platt, S. M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa,
645 M., Ciarelli, G., Piazzalunga, A., Schwikowski, M., Abbaszade, G., Schnelle-Kreis, J.,
646 Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., Haddad, I. E., and Prevot, A. S. H.: High
647 secondary aerosol contribution to particulate pollution during haze events in China, *Nature*, 514,
648 218–222, 2014.
- 649 Huang, R. J., He, Y., Duan, J., Li, Y., Chen, Q., Zheng, Y., Chen, Y., Hu, W., Lin, C., Ni, H.,
650 Dai, W., Cao, J., Wu, Y., Zhang, R., Xu, W., Ovadnevaite, J., Ceburnis, D., Hoffmann, T., and
651 O'Dowd, C. D.: Contrasting sources and processes of particulate species in haze days with low
652 and high relative humidity in wintertime Beijing, *Atmos. Chem. Phys.*, 20, 9101–9114,
653 10.5194/acp-20-9101-2020, 2020.
- 654 Jang, M., and Kamens, R. M.: Atmospheric secondary aerosol formation by heterogeneous
655 reactions of aldehydes in the presence of a sulfuric acid aerosol catalyst, *Environ. Sci. Technol.*,
656 35, 4758–4766, 10.1021/es010790s, 2001.
- 657 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H.,
658 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M.,
659 Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin,
660 C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M.,
661 Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, J., Huffman, J. A.,
662 Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick,



- 663 F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A.,
664 Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel,
665 J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C.,
666 Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of organic
667 aerosols in the atmosphere, *Science*, 326, 1525–1529, <https://doi.org/10.1126/science.1180353>,
668 2009.
- 669 Kautzman, K., Surratt, J., Chan, M., Chan, A., Hersey, S., Chhabra, P., Dalleska, N., Wennberg,
670 P., Flagan, R., and Seinfeld, J.: Chemical composition of gas-and aerosol-phase products from
671 the photooxidation of naphthalene, *J. Phys. Chem. A*, 114, 913–934, 2009.
- 672 Knipping, E. M., Lakin, M. J., Foster, K. L., Jungwirth, P., Tobias, D. J., Gerber, R. B., Dabdub,
673 D., and Finlayson-Pitts, B. J.: Experiments and simulations of ion-enhanced interfacial chemistry
674 on aqueous NaCl aerosols, *Science*, 288, 301, [10.1126/science.288.5464.301](https://doi.org/10.1126/science.288.5464.301), 2000.
- 675 Krapf, M., El Haddad, I., Bruns, E. A., Molteni, U., Daellenbach, K. R., Prévôt, A. S.,
676 Baltensperger, U., and Dommen, J.: Labile peroxides in secondary organic aerosol, *Chem*, 1,
677 603–616, 2016.
- 678 Laskin, A., Gaspar, D. J., Wang, W., Hunt, S. W., Cowin, J. P., Colson, S. D., and Finlayson-
679 Pitts, B. J.: Reactions at interfaces as a source of sulfate formation in sea-salt particles, *Science*,
680 301, 340, [10.1126/science.1085374](https://doi.org/10.1126/science.1085374), 2003.
- 681 Leng, C. B., Roberts, J. E., Zeng, G., Zhang, Y. H., and Liu, Y.: Effects of temperature, pH, and
682 ionic strength on the Henry's law constant of triethylamine, *Geophys. Res. Lett.*, 42, 3569–3575,
683 [10.1002/2015gl063840](https://doi.org/10.1002/2015gl063840), 2015.
- 684 Li, G., Bei, N., Cao, J., Huang, R., Wu, J., Feng, T., Wang, Y., Liu, S., Zhang, Q., Tie, X., and
685 Molina, L. T.: A possible pathway for rapid growth of sulfate during haze days in China, *Atmos.*
686 *Chem. Phys.*, 17, 3301–3316, <https://doi.org/10.5194/acp17-3301-2017>, 2017.
- 687 Li, J., Zhang, Y., Cao, F., Zhang, W., Fan, M., Lee, X., Michalski, G.: Stable sulfur isotopes
688 revealed a major role of transition-metal-ion catalyzed SO₂ oxidation in haze episodes,
689 [doi:10.1021/acs.est.9b07150](https://doi.org/10.1021/acs.est.9b07150), 2020.
- 690 Lind, J. A., Lazrus, A. L., and Kok, G. L.: Aqueous phase oxidation of sulfur (IV) by hydrogen
691 peroxide, methylhydroperoxide, and peroxyacetic acid, *J. Geophys. Res. Atmos.*, 92, 4171–4177,
692 1987.
- 693 Liu, C., Chen, T., Liu, Y., Liu, J., He, H., Zhang, P.: Enhancement of secondary organic aerosol
694 formation and its oxidation state by SO₂ during photooxidation of 2-methoxyphenol, *Atmos.*
695 *Chem. Phys.*, 19, 2687–2700, 2019.
- 696 Liu, L., Bei, N., Wu, J., Liu, S., Zhou, J., Li, X., Yang, Q., Feng, T., Cao, J., Tie, X. and Li, G.:
697 Effects of stabilized Criegee intermediates (sCIs) on sulfate formation: a sensitivity analysis
698 during summertime in Beijing–Tianjin–Hebei (BTH), China, *Atmos. Chem. Phys.*, 19, 13341–
699 13354, 2019.
- 700 Liu, T., Clegg, S. L. and Abbatt, J. P. D.: Fast oxidation of sulfur dioxide by hydrogen peroxide
701 in deliquesced aerosol particles, *Proc. Natl. Acad. Sci. U. S. A.*, 117, 1354–1359, 2020.
- 702 Liu, Y., Liggio, J., Staebler, R., and Li, S. M.: Reactive uptake of ammonia to secondary organic
703 aerosols: kinetics of organonitrogen formation, *Atmos. Chem. Phys.*, 15, 13569–13584,
704 [doi:10.5194/acp-15-13569-2015](https://doi.org/10.5194/acp-15-13569-2015), 2015.
- 705 Maaß, F., Elias, H., and Wannowius, K. J.: Kinetics of the oxidation of hydrogen sulfite by
706 hydrogen peroxide in aqueous solution: ionic strength effects and temperature dependence,
707 *Atmos. Environ.*, 33, 4413–4419, [https://doi.org/10.1016/S1352-2310\(99\)00212-5](https://doi.org/10.1016/S1352-2310(99)00212-5), 1999.



- 708 Mauldin, R. L., Berndt, T., Sipilä, M., Paasonen, P., Petäjä, T., Kim, S., Kurtén, T., Stratmann,
709 F., Kerminen, V. M., and Kulmala, M.: A new atmospherically relevant oxidant of sulphur
710 dioxide, *Nature*, 488, 193–196, <https://doi.org/10.1038/nature11278>, 2012.
- 711 Mekic, M., Loisel, G., Zhou, W., Jiang, B., Vione, D., and Gligorovski, S.: Ionic-strength effects
712 on the reactive uptake of ozone on aqueous pyruvic acid: Implications for air–sea ozone
713 deposition, *Environ. Sci. Technol.*, 52, 12306–12315, [10.1021/acs.est.8b03196](https://doi.org/10.1021/acs.est.8b03196), 2018.
- 714 Mekic, M., Zeng, J., Zhou, W., Loisel, G., Jin, B., Li, X., Vione, D., and Gligorovski, S.: Ionic
715 strength effect on photochemistry of fluorene and dimethylsulfoxide at the air–sea interface:
716 Alternative formation pathway of organic sulfur compounds in a marine atmosphere, *ACS Earth
717 Space Chem.*, 4, 1029–1038, [10.1021/acsearthspacechem.0c00059](https://doi.org/10.1021/acsearthspacechem.0c00059), 2020.
- 718 Mishra, H., Enami, S., Nielsen, R. J., Hoffmann, M. R., Goddard, W. A., and Colussi, A. J.:
719 Anions dramatically enhance proton transfer through aqueous interfaces, *Proc. Natl. Acad. Sci.
720 U. S. A.*, 109, 10228–10232, 2012.
- 721 Ng, N., Kroll, J., Chan, A., Chhabra, P., Flagan, R., and Seinfeld, J.: Secondary organic aerosol
722 formation from m-xylene, toluene, and benzene, *Atmos. Chem. Phys.*, 7, 3909–3922,
723 <http://www.atmos-chem-phys.net/7/3909/2007/>, 2007.
- 724 Ng, N. L., Kwan, A. J., Surratt, J. D., Chan, A. W. H., Chhabra, P. S., Sorooshian, A., Pye, H. O.
725 T., Crounse, J. D., Wennberg, P. O., Flagan, R. C., and Seinfeld, J. H.: Secondary organic
726 aerosol (SOA) formation from reaction of isoprene with nitrate radicals (NO₃), *Atmos. Chem.
727 Phys.*, 8, 4117–4140, <http://www.atmos-chem-phys.net/8/4117/2008/>, 2008.
- 728 O’Brien, R. E., Wang, B., Kelly, S. T., Lundt, N., You, Y., Bertram, A. K., Leone, S. R., Laskin,
729 A., and Gilles, M. K.: Liquid–liquid phase separation in aerosol particles: Imaging at the
730 nanometer scale, *Environ. Sci. Technol.*, 49, 4995–5002, [10.1021/acs.est.5b00062](https://doi.org/10.1021/acs.est.5b00062), 2015.
- 731 Pankow, J. F. and Asher, W. E.: SIMPOL.1: a simple group contribution method for predicting
732 vapor pressures and enthalpies of vaporization of multifunctional organic compounds, *Atmos.
733 Chem. Phys.*, 8, 2773–2796, <http://www.atmos-chem-phys.net/8/2773/2008/>, 2008.
- 734 Qiu, J., Liang, Z., Tonokura, K., Colussi, A. J., and Enami, S.: Stability of monoterpene-derived
735 α -hydroxyalkyl-hydroperoxides in aqueous organic media – relevance to the fate of
736 hydroperoxides in aerosol particle phases, *Environ. Sci. Technol.*, [10.1021/acs.est.9b07497](https://doi.org/10.1021/acs.est.9b07497),
737 2020.
- 738 Rodríguez-Sevilla, J., Álvarez, M., Limiñana, G., Díaz, M. C.: Dilute SO₂ absorption equilibria
739 in aqueous HCl and NaCl solutions at 298.15 K, *J. Chem. Eng. Data*, 47, 1339–1345, 2002.
- 740 Ruiz-Lopez, M.F., Francisco, J.S., Martins-Costa, M.T. and Anglada, J.M.: Molecular reactions
741 at aqueous interfaces. *Nat. Rev. Chem.*, 1–17, 2020.
- 742 Seinfeld, J. H., and Pandis, S. N.: *Atmospheric chemistry and physics: from air pollution to
743 climate change*, John Wiley & Sons, 2012.
- 744 Sha, T., Ma, X., Jia, H., Tian, R., Chang, Y., Cao, F., and Zhang, Y.: Aerosol chemical
745 component: Simulations with WRF-Chem and comparison with observations in Nanjing, *Atmos.
746 Environ.*, 218, 116982, <https://doi.org/10.1016/j.atmosenv.2019.116982>, 2019.
- 747 Shang, J., Li, J., Zhu, T.: Heterogeneous reaction of SO₂ on TiO₂ particles. *Sci.
748 China Chem.*, 53, 2637–2643, 2010.
- 749 Shi, Q., Davidovits, P., Jayne, J. T., Worsnop, D. R., and Kolb, C. E.: Uptake of gas-phase
750 ammonia. 1. Uptake by aqueous surfaces as a function of pH, *J. Phys. Chem. A*, 103, 8812–8823,
751 [10.1021/jp991696p](https://doi.org/10.1021/jp991696p), 1999.
- 752 Song, M., Marcolli, C., Krieger, U. K., Zuend, A., and Peter, T.: Liquid–liquid phase separation



- 753 in aerosol particles: dependence on O:C, organic functionalities, and compositional complexity,
754 *Geophys. Res. Lett.*, 39, L19801, doi:10.1029/2012GL052807, 2012.
- 755 Song, S., Gao, M., Xu, W., Shao, J., Shi, G., Wang, S., Wang, Y., Sun, Y., and McElroy, M. B.:
756 Fine-particle pH for Beijing winter haze as inferred from different thermodynamic equilibrium
757 models, *Atmos. Chem. Phys.*, 18, 7423–7438, <https://doi.org/10.5194/acp-18-7423-2018>, 2018.
- 758 Song, S., Gao, M., Xu, W., Sun, Y., Worsnop, D. R., Jayne, J. T., Zhang, Y., Zhu, L., Li, M.,
759 Zhou, Z., Cheng, C., Lv, Y., Wang, Y., Peng, W., Xu, X., Lin, N., Wang, Y., Wang, S., Munger,
760 J. W., Jacob, D. J., and McElroy, M. B.: Possible heterogeneous chemistry of hydroxy
761 methanesulfonate (HMS) in northern China winter haze, *Atmos. Chem. Phys.*, 19, 1357–1371,
762 <https://doi.org/10.5194/acp-19-1357-2019>, 2019.
- 763 Su, H., Cheng, Y., and Poschl, U.: New multiphase chemical processes influencing atmospheric
764 aerosols, air quality, and climate in the anthropocene, *Acc. Chem. Res.*, e1601530-2983,
765 <https://doi.org/10.1021/acs.accounts.0c00246>, 2020.
- 766 Sun, Y., Wang, Z., Fu, P., Jiang, Q., Yang, T., Li, J., and Ge, X.: The impact of relative humidity
767 on aerosol composition and evolution processes during wintertime in Beijing, China, *Atmos.*
768 *Environ.*, 77, 927-934, 2013.
- 769 Surratt, J. D., Murphy, S. M., Kroll, J. H., Ng, N. L., Hildebrandt, L., Sorooshian, A.,
770 Szmigielski, R., Vermeylen, R., Maenhaut, W., and Claeys, M.: Chemical composition of
771 secondary organic aerosol formed from the photooxidation of isoprene, *J. Phys. Chem. A*, 110,
772 9665-9690, 2006.
- 773 Thornton, J. A., Braban, C. F., and Abbatt, J. P. D.: N₂O₅ hydrolysis on sub-micron organic
774 aerosols: the effect of relative humidity, particle phase, and particle size, *Phys. Chem. Chem*
775 *Phys.*, 5, 4593–4603, <https://doi.org/10.1039/B307498F>, 2003.
- 776 Tie, X., Brasseur, G., Emmons, L., Horowitz, I., and Kinnison, D.: Effects of aerosols on
777 tropospheric oxidants: a global model study, *J. Geophys. Res. Atmos.*, 106, 22931–22964, 2001.
- 778 Usher, C. R., Al-Hosney, H., Carlos-Cuellar, S., and Grassian, V. H.: A laboratory study of the
779 heterogeneous uptake and oxidation of sulfur dioxide on mineral dust particles, *J. Geophys.*
780 *Res.*, 107, 4713, doi:10.1029/2002JD002051, 2002.
- 781 Varutbangkul, V., Brechtel, F. J., Bahreini, R., Ng, N. L., Keywood, M. D., Kroll, J. H., Flagan,
782 R. C., Seinfeld, J. H., Lee, A., and Goldstein, A. H.: Hygroscopicity of secondary organic
783 aerosols formed by oxidation of cycloalkenes, monoterpenes, sesquiterpenes, and related
784 compounds, *Atmos. Chem. Phys.*, 6, 2367–2388, [http://www.atmos-chem-phys.net/6/23](http://www.atmos-chem-phys.net/6/2367/2006/)
785 [67/2006/](http://www.atmos-chem-phys.net/6/2367/2006/), 2006.
- 786 Veghte, D. P., Altaf, M. B., and Freedman, M. A.: Size dependence of the structure of organic
787 aerosol, *J. Am. Chem. Soc.*, 135, 16046–16049, 2013.
- 788 Wang, G., Zhang, R., Gomez, M. E., Yang, L., Zamora, M. L., Hu, M., Lin, Y., Peng, J., Guo, S.,
789 and Meng, J.: Persistent sulfate formation from London Fog to Chinese haze, *Proc. Natl. Acad.*
790 *Sci. U. S. A.*, 113, 13630-13635, 2016.
- 791 Wang, S., Ye, J., Soong, R., Wu, B., Yu, L., Simpson, A. J., and Chan, A. W. H.: Relationship
792 between chemical composition and oxidative potential of secondary organic aerosol from
793 polycyclic aromatic hydrocarbons, *Atmos. Chem. Phys.*, 18, 3987-4003, 2018.
- 794 Wang, S., Zhou, S., Tao, Y., Tsui, W. G., Ye, J., Yu, J. Z., Murphy, J. G., McNeill, V. F.,
795 Abbatt, J. P. D., and Chan, A. W. H.: Organic peroxides and sulfur dioxide in aerosol: Source of
796 particulate sulfate, *Environ. Sci. Technol.*, 53, 10695-10704, [10.1021/acs.est.9b02591](https://doi.org/10.1021/acs.est.9b02591), 2019.
- 797 Wang, X.; Gemayel, R.; Hayeck, N.; Perrier, S.; Charbonnel, N.; Xu, C.; Chen, H.; Zhu, C.;
798 Zhang, L.; Wang, L.; Nizkorodov, S. A.; Wang, X.; Wang, Z.; Wang, T.; Mellouki, A.; Riva, M.;



- 799 Chen, J.; George, C. Atmospheric photosensitization: A new pathway for sulfate formation,
800 *Environ. Sci. Technol.*, 54, 3114–3120, 2020.
- 801 Wang, Y., Zhang, Q., Jiang, J., Zhou, W., Wang, B., He, K., Duan, F., Zhang, Q., Philip, S., and
802 Xie, Y.: Enhanced sulfate formation during China’s severe winter haze episode in January 2013
803 missing from current models, *J. Geophys. Res. Atmos.*, 119, 10425–10440, <https://doi.org/10.1029/2013JD021426>, 2014.
- 804
805 Wei, H., Vejerano, E. P., Leng, W., Huang, Q., Willner, M. R., Marr, L. C., and Vikesland, P. J.:
806 Aerosol microdroplets exhibit a stable pH gradient, *Proc. Natl. Acad. Sci. U. S. A.*, 115, 7272,
807 [10.1073/pnas.1720488115](https://doi.org/10.1073/pnas.1720488115), 2018.
- 808 Xu, L., Guo, H., Boyd, C. M., Klein, M., Bougiatioti, A., Cerully, K. M., Hite, J. R., Isaacman-
809 VanWertz, G., Kreisberg, N. M., and Knote, C.: Effects of anthropogenic emissions on aerosol
810 formation from isoprene and monoterpenes in the southeastern United States, *Proc. Natl. Acad. Sci. U. S. A.*, 112, 37–42, 2015.
- 811
812 Yang, Y., Wang, H., Smith, S. J., Easter, R., Ma, P.-L., Qian, Y., Yu, H., Li, C., and Rasch, P. J.:
813 Global source attribution of sulfate concentration and direct and indirect radiative forcing,
814 *Atmos. Chem. Phys.*, 17, 8903–8922, <https://doi.org/10.5194/acp17-8903-2017>, 2017.
- 815 Yao, M., Zhao, Y., Hu, M., Huang, D., Wang, Y. C., Yu, J. Z., and Yan, N.: Multiphase reactions
816 between secondary organic aerosol and sulfur dioxide: kinetics and contributions to sulfate
817 formation and aerosol aging, *Environ. Sci. Tech. Lett.* 6, 768–774, [10.1021/acs.estlett.9b00657](https://doi.org/10.1021/acs.estlett.9b00657), 2019.
- 818
819 Ye, J., Gordon, C. A., and Chan, A. W. H.: Enhancement in secondary organic aerosol formation
820 in the presence of preexisting organic particle, *Environ. Sci. Technol.*, 50, 3572–3579, 2016.
- 821 Ye, J., Abbatt, J. P. D., and Chan, A. W. H.: Novel pathway of SO₂ oxidation in the atmosphere:
822 reactions with monoterpene ozonolysis intermediates and secondary organic aerosol, *Atmos. Chem. Phys.*, 18, 5549–5565, <https://doi.org/10.5194/acp18-5549-2018>, 2018.
- 823
824 Yee, L. D., Isaacman-VanWertz, G., Wernis, R. A., Kreisberg, N. M., Glasius, M., Riva, M.,
825 Surratt, J. D., de Sá, S. S., Martin, S. T., Alexander, M. L., Palm, B. B., Hu, W., Campuzano-
826 Jost, P., Day, D. A., Jimenez, J. L., Liu, Y., Misztal, P. K., Artaxo, P., Viegas, J., Manzi, A., de
827 Souza, R. A. F., Edgerton, E. S., Baumann, K., and Goldstein, A. H.: Natural and
828 anthropogenically influenced isoprene oxidation in southeastern United States and central
829 Amazon, *Environ. Sci. Technol.*, 54, 5980–5991, [10.1021/acs.est.0c00805](https://doi.org/10.1021/acs.est.0c00805), 2020.
- 830 You, Y., Renbaum-Wolff, L., Bertram, A. K.: Liquid-liquid phase separation in particles
831 containing organics mixed with ammonium sulfate, ammonium bisulfate, ammonium nitrate or
832 sodium chloride, *Atmos. Chem. Phys.*, 13, 11723–11734, <https://doi.org/10.5194/acp-13-11723-2013>, 2013.
- 833
834 You, Y., Smith, M. L., Song, M., Martin, S. T., and Bertram, A. K.: Liquid–liquid phase
835 separation in atmospherically relevant particles consisting of organic species and inorganic salts,
836 *Int. Rev. Phys. Chem.*, 33, 43–77, [doi:10.1080/0144235X.2014.890786](https://doi.org/10.1080/0144235X.2014.890786), 2014.
- 837 Zhang, S., Xing, J., Sarwar, G., Ge, Y., He, H., Duan, F., Zhao, Y., He, K., Zhu, L. and Chu, B.:
838 Parameterization of heterogeneous reaction of SO₂ to sulfate on dust with coexistence of NH₃
839 and NO₂ under different humidity conditions, *Atmos. Environ.*, 208, 133–140, 2019.
- 840 Zhao, Y., Liu, Y., Ma, J., Ma, Q., and He, H.: Heterogeneous reaction of SO₂ with soot: The
841 roles of relative humidity and surface composition of soot in surface sulfate formation, *Atmos. Environ.*, 152, 465–476, 2017.
- 842
843 Zheng, B., Zhang, Q., Zhang, Y., He, K. B., Wang, K., Zheng, G. J., Duan, F. K., Ma, Y. L., and
844 Kimoto, T.: Heterogeneous chemistry: a mechanism missing in current models to explain



845 secondary inorganic aerosol formation during the January 2013 haze episode in North China,
846 Atmos. Chem. Phys., 15, 2031–2049, doi:10.5194/acp-15-2031-2015, 2015.
847 Zheng, G. J., Duan, F. K., Su, H., Ma, Y. L., Cheng, Y., Zheng, B., Zhang, Q., Huang, T.,
848 Kimoto, T., Chang, D., Pöschl, U., Cheng, Y. F., and He, K. B.: Exploring the severe winter haze
849 in Beijing: the impact of synoptic weather, regional transport and heterogeneous reactions,
850 Atmos. Chem. Phys., 15, 2969–2983, doi:10.5194/acp-15-2969-2015, 2015.

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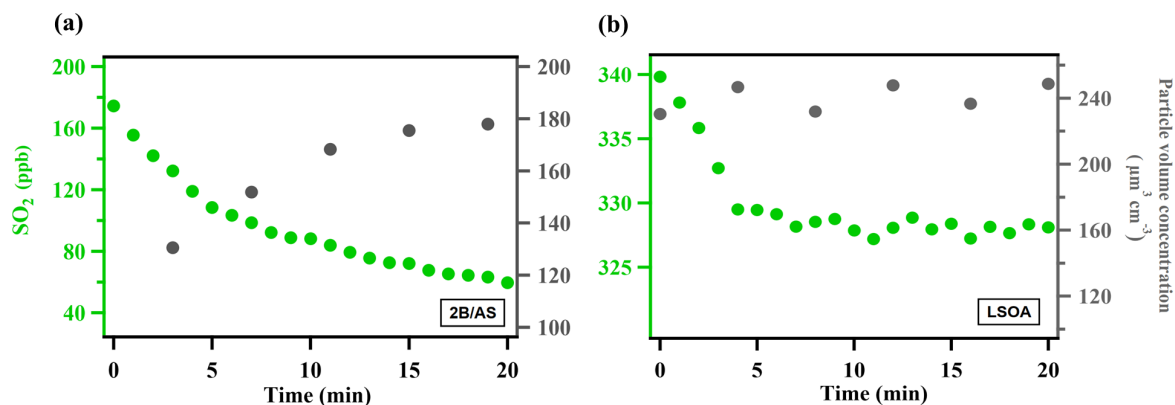
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875 **Figure 1.** Typical evolution of the species monitored during γ_{SO_2} measurement for (a)
876 ammonium sulfate mixed with 2-butanone organic peroxide (2B/AS, Expt. 16) and (b) limonene
877 SOA (LSOA, Expt. 27). Particle volume concentrations measured by SMPS have been corrected
878 for wall loss assuming a pseudo first-order loss rate (Ye et al., 2016). γ_{SO_2} was calculated for the
879 initial portion of the decay (first 7 minutes).

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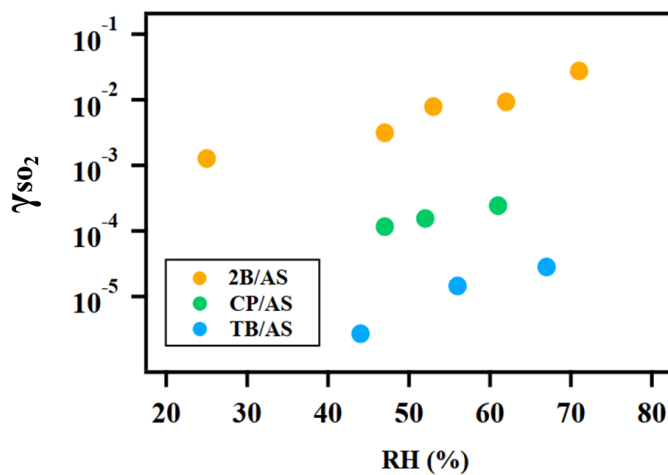
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887 **Figure 2.** Exponential relationship between γ_{SO_2} and RH for ammonium sulfate
888 containing 2-butanone peroxide (2B), cumene hydroperoxide (CP), tert-butyl hydroperoxide
889 (TB).

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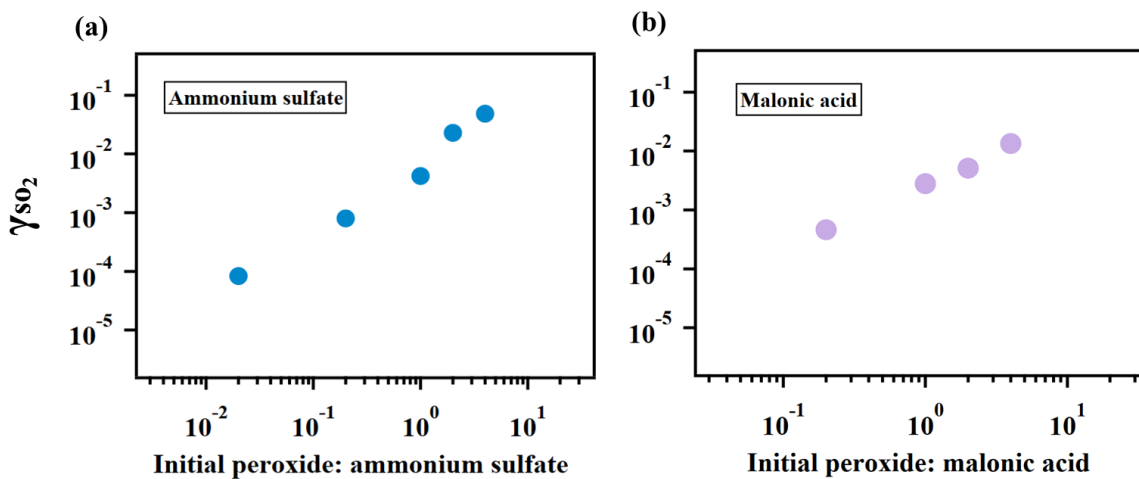
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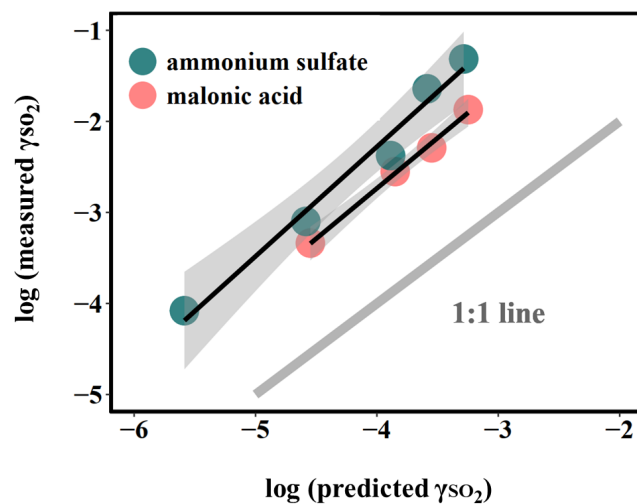
899 **Figure 3.** Relationship between γ_{SO_2} and particulate peroxide content. γ_{SO_2} for ammonium sulfate
900 (a) and malonic acid aerosol (b) containing different amount of 2-butanone peroxide are shown
901 here. The observed dependence of γ_{SO_2} on the amount of peroxide injected are linear since the
902 slopes of the relationship are both nearly 1 in (a) and (b).

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908 **Figure 4.** Relationship between measured γ_{SO_2} and γ_{SO_2} predicted by Eqn. 4. The large deviation
909 from the 1:1 line, which represents the difference between the measured uptake coefficient and
910 predicted values based on kinetics in the dilute aqueous phase, indicates that aerosol reactive
911 uptake is significantly faster than reactions in dilute aqueous phase. This enhancement is likely
912 driven in part by high ionic strengths, as the difference between measured γ_{SO_2} and predicted γ_{SO_2}
913 are consistently higher for organic peroxide containing ammonium sulfate (high ionic strength)
914 than for that mixed with malonic acid (lower ionic strength).

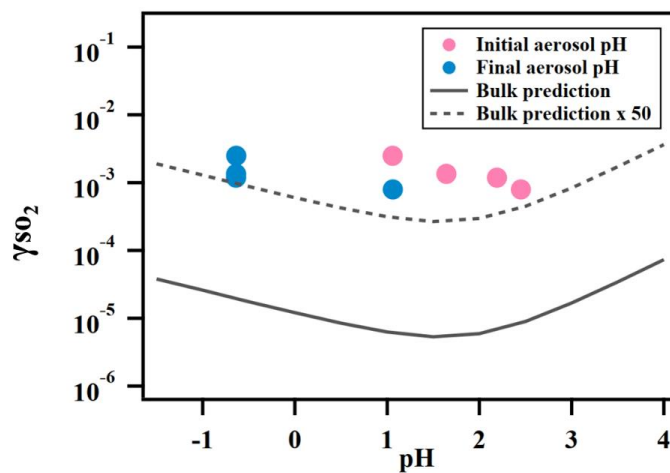
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921 **Figure 5.** Relationship between γ_{SO_2} and aerosol phase pH for ammonium sulfate aerosol
922 containing 2-butanone peroxide.

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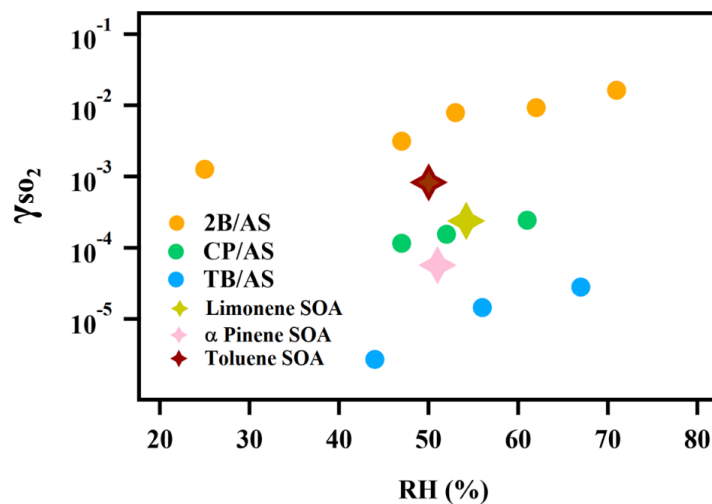
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932 **Figure 6.** γ_{SO_2} measured for different types of organic aerosol. The reactive uptake coefficient of
933 SO_2 onto SOA are on the order of 10^{-4} .

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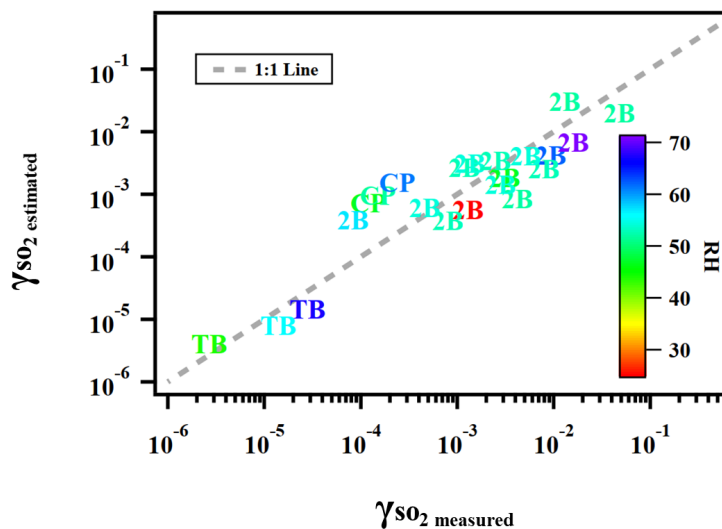
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944 **Figure 7.** Predicted γ_{SO_2} using Equation (8) versus measured γ_{SO_2} for ammonium sulfate or
945 malonic acid aerosol containing 2-butanone peroxide (2B), cumene hydroperoxide (CP), tert-
946 butyl hydroperoxide (TB) under different experimental conditions.

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