



1 Turbulence-permitting air pollution simulation for the

2 Stuttgart metropolitan area

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Abstract. Air pollution is one of the major challenges in urban areas. It can have a major impact on human health and society and is currently a subject of several litigations at European courts. Information on the level of air pollution is based on near surface measurements, which are often irregularly distributed along the main traffic roads and provide almost no information about the residential areas and office districts in the cities. To further enhance the process understanding and give scientific support to decision makers, we developed a prototype for an air quality forecasting system (AQFS) within the EU demonstration project "Open Forecast".

For AQFS, the Weather Research and Forecasting model together with its coupled chemistry component (WRF-Chem) is applied for the Stuttgart metropolitan area in Germany. Three model domains from 1.25 km down to a turbulence permitting resolution of 50 m were used and a single layer urban canopy model was active in all domains. As demonstration case study the 21 January 2019 was selected which was a heavy polluted day with observed PM_{10} concentrations exceeding 50 µg m⁻³.

Our results show that the model is capable to reasonably simulate the diurnal cycle of surface fluxes and 2-m temperatures as well as evolution of the stable and shallow boundary layer typically occurring in wintertime in Stuttgart. The simulated fields of particulates with a diameter of less than 10 μ m (PM₁₀) and Nitrogen dioxide (NO₂) allow a clear statement about the most heavily polluted areas apart from the irregularly distributed measurement sites. Together with information about the vertical distribution of PM₁₀ and NO₂ from the model, AQFS will serve as a valuable tool for air quality forecast and has the potential of being applied to other cities around the world.

28 1. Introduction

Currently more than 50 % of the global population live in cities whereas the United Nations (UN) expect a further
 increase by about 10 % in 2030 (UN, 2018). The UN also expect that in 2030 34% of the world population will
 reside in cities with more than 500 000 inhabitants.

32 To protect human life, the World Health Organization (WHO) proposed maximum permittable pollution levels

 $\textbf{33} \qquad (Maynard \ et \ al., \ 2017 \ and \ references \ therein). \ E.g. \ for \ particulate \ matter \ with \ particle \ diameters \ less \ than \ 10 \ \mu m$

34 (PM₁₀), the critical value is an annual mean concentration of 20 μ g m⁻³ or a daily mean value of 50 μ g m⁻³ (WHO,

- 35 2005). For Nitrogen dioxide (NO₂) the critical values are 200 μ g m⁻³ and 40 μ g m⁻³ as daily and annual mean
- 36 values, respectively.





37 Due to a strong increase of road traffic in major European cities (Thunis et al., 2017), these pollution levels are 38 often violated in larger cities. This can lead to health and environmental problems and is currently part of several 39 litigations e.g. at the German Federal Administrative Court dealing with possible driving bans for non low-40 emission vehicles. The basis for these litigations are mostly few local observations which are unevenly distributed. 41 In combination with special meteorological conditions like winter time thermal inversion layers it can be 42 misleading to conclude about the overall air quality in the city. According to e.g. the German Federal Immission 43 Control Ordinance¹ it is sufficient that traffic related measurements are representative for a section of 100 m, but 44 this is not representative for the commercial and office districts in the cities that are suffering from traffic control 45 in case of fine dust alerts and residential areas. Namely in residential areas health protection action plans require 46 representative air quality measures.

47 Therefore it becomes important to apply a more scientifically valid approach by applying coupled atmospheric 48 and chemistry models to predict air quality. Regional and global atmospheric models like the Weather Research 49 and Forecasting (WRF) model (Skamarock et al., 2019), the Consortium for Small Scale modeling (COSMO; 50 Baldauf et al., 2011), the Icosahedric Nonhydrostatic model (ICON; Zängl et al., 2015), or the Regional Climate 51 Model system (RegCM4; Giorgi et al., 2012) are often used to force offline chemistry transport models like 52 CHIMERE (Mailler et al., 2017), LOTOS-EUROS (Manders et al., 2017), EURopean Air Pollution Dispersion 53 (EURAD; Memmesheimer et al., 2004), and Model for OZone And Related chemical Tracers (MOZART) 54 (Brasseur et al., 1998; Horowitz et al., 2003).

Several studies showed that combining an atmospheric model with an online coupled chemistry component is a
suitable tool for air quality and pollution modeling in urban areas at the convection permitting (CP) resolution
(Fallmann et al., 2014; Kuik et al., 2016; Zhong et al., 2016; Kuik et al., 2018; Huszar et al., 2020).

Compared to chemical transport models, coupled models like WRF-Chem (Grell et al., 2005), COSMO-ART
(Vogel et al., 2009), ICON-ART (Rieger et al., 2015), and the Integrated Forecasting System (IFS) MOZART
(Flemming et al., 2015) allow for a direct interaction of aerosols with radiation leading to a better representation
of the energy balance closure at the surface as it would be the case when applying an offline chemistry model.

62 As usually the terrain and land cover over urban areas show fine scale structures which are not resolved even by a 63 CP resolution, there is a need for turbulence permitting (TP) simulations with horizontal grid increments of a few 64 hundred meters or even less. Important features are, e.g., urban heat island effects (Fallmann et al., 2014; Fallmann 65 et al., 2016; García-Díez et al., 2016; Li et al., 2019) and local wind systems like mountain and valley winds due 66 to differential heating (Corsmeier et al., 2011; e.g. Jin et al., 2016). Also, micro- and mesoscale wind systems can 67 develop due to urban structures and the heterogeneity of the land surface. It is well known that TP simulations are 68 a promising tool to further enhance the understanding of processes in the atmospheric boundary layer (Heinze et al., 2017b; Panosetti et al., 2016; Heinze et al., 2017a; Bauer et al., 2020) in urban areas (Nakayama et al., 2012; 69 70 Maronga et al., 2019; Maronga et al., 2020).

In order to further enhance the quality of the simulations, building and urban canopy models (UCM) are developed
 (Martilli et al., 2002; Kusaka and Kimura, 2004; Salamanca and Martilli, 2010; Maronga et al., 2019; Scherer et
 al., 2019; Teixeira et al., 2019). The main purpose of UCMs is to provide a better description of the lower

¹<u>https://www.gesetze-im-internet.de/bimschv_39/anlage_3.html</u>





74 boundaries over urban areas such as building, roof and road geometries and their interactions with atmospheric

- 75 water vapor, wind, and radiation.
- With the EU-funded project Open Forecast (https://open-forecast.eu/en/) it was intended to develop a prototype for an air quality forecasting system (AQFS) for the Stuttgart metropolitan area in southwest Germany. Open Forecast is a demonstration project to show the potential of open data combined with supercomputer resources to create new data products for European citizens and public authorities. The long-term goal is to provide end users and political decision-makers a useful tool, particularly considering further urbanization, heat island effects as well as potential driving restrictions due to recent EU decisions on emission limits.
- 82 For our AQFS we use the WRF-Chem NWP model (Grell et al., 2005; Skamarock et al., 2019) as the WRF model 83 is extensively evaluated over Europe at different time scales and horizontal resolutions (San José et al., 2013; 84 Warrach-Sagi et al., 2013; Milovac et al., 2016; Lian et al., 2018; Molnár et al., 2019; Bauer et al., 2020; Coppola 85 et al., 2020; Schwitalla et al., 2020). It can easily be set up in a nested configuration over all regions of the Earth. 86 Compared to PALM-4U model, the nested model domains are driven by the full atmospheric and chemical 87 information from the parent domain along its lateral boundaries. Also, it contains well-characterized combinations 88 of parameterizations of turbulence and cloud microphysics in the outer domain that are consistent with the inner 89 TP domains where the high-quality cloud parameterization remains. No switch between different model systems 90 is required, which is expected to provide a great advantage with respect to the skill of air pollution and 91 meteorological forecasts.

92 To enhance the forecast skill, suitable variational and ensemble-based data assimilation systems are already in
93 place to further improve the meteorological initial conditions (Barker et al., 2012; Zhang et al., 2014; Kawabata et
94 al., 2018; Thundathil et al., 2020) and the chemical initial conditions (Chen et al., 2019; Sun et al., 2020) but this
95 is beyond the scope of our study.

- 96 The Parallelized Large-Eddy Simulation Model (PALM) model (Maronga et al., 2015) is another widely used TP
 97 simulation model over Europe. PALM did not include the full interaction between land-surface, radiation, cloud
 98 microphysics and chemistry during the performance of our study. The very recent version 6.0 of PALM-4U
 99 (PALM for urban applications) (Maronga et al., 2020) is expected to contain a fully coupled chemistry module
 100 (Khan et al., 2020, under review).
- 101 Fallmann et al. (2016) and Kuik et al. (2016) performed air quality simulations with WRF-Chem over the cities of 102 Berlin and Stuttgart on a CP resolution down to 1km and less than 40 model levels. They used the TNO-MACC 103 emission inventory (Kuenen et al., 2014) which is available as an annual totals on a 7 km x 7 km resolution. As 104 the topography of Stuttgart is very complex, the AQFS applies the WRF-Chem model on a turbulence permitting 105 horizontal resolution using 100 model levels to account for the shallow boundary layer occurring during 106 wintertime. In addition, we applied a local emission data set from the Baden-Württemberg State Institute for the 107 Environment, Survey and Nature Conservation available as annual mean on a horizontal resolution of 500 m x 500 108 m to resolve fine-scale emission structures.
- Our study focuses on the methodology how to set up a AQFS prototype by using WRF-Chem and its applicationto a typical wintertime situation in the Stuttgart metropolitan area. The manuscript is set up as follows: section 2
- 111 describes the design of our AQFS model system on the turbulence permitting resolution of 50 m followed by a





- 112 description of the selected case study. Section 4 shows the results including a discussion, sect. 5 summarizes our
- 113 work and gives an outlook on potential future enhancements of the AQFS prototype.

114 2. AQFS design

115 2.1. WRF model set-up

For our AQFS, we selected the Advanced Research WRF-Chem model in version 4.0.3 (Grell et al., 2005; Skamarock et al., 2019). To reach the targeted resolution of 50 m, three model domains have been applied with horizontal resolutions of 1250 m, 250 m, and 50 m and encompasses 800*800 grid cells in the outer domain and 601*601 grid cells in the two inner TP domains. The reason to start with a resolution of 1250 m in the outermost domain is to avoid the application of a convection parametrization which can deteriorate the model results (Prein et al., 2015; Coppola et al., 2020). The areas of model domain 1 and 3 are shown in Fig. 1.

As seen from Fig. 1b, the Stuttgart metropolitan area is characterized by an elevation variation of more than 300 m. The lowest elevation is approx. 220 m in the basin while the highest elevation reaches up to 570 m. As the main traffic roads are in the basin, especially during wintertime this often leads to a worsening of the air quality as the surrounding prevents an air mass exchange due to the stationary temperature inversion.

126 For the WRF model system land cover and soil texture fields are not available at resolutions higher than 500m. 127 Therefore we reclassified land cover data from the Copernicus CLC 2012 data set (European Union, 2012), 128 available on a resolution of 100 m, from the original 44 categories to the categories applied in the WRF model for 129 the simulations of the outer 2 domains. For the innermost model domain, we incorporated the most recent high-130 resolution land-cover data set from the Baden-Württemberg State Institute for the Environment (LUBW), which 131 is derived from Landsat (Butcher et al., 2019) in 2010 and is available at 30 m resolution (https://udo.lubw.baden-132 wuerttemberg.de/public/) This data set was also reclassified to the corresponding land cover categories used in 133 WRF and is shown in Fig. 2.

134 The resolution of the provided default Food and Agriculture Organization of the United Nations (FAO) soil texture 135 data is only 10 km, therefore we used soil texture data from the International Soil Reference and Information Centre (ISRIC) SoilGrids project (Hengl et al., 2014; Hengl et al., 2015). These data are available on a resolution 136 137 of 250 m. Terrain information was provided by the National Center for Atmospheric Research (NCAR) derived 138 from the Global multi-resolution terrain elevation data 2010 (GMTED2010) data set (Danielson and Gesch, 2011) for domain 1. As the horizontal resolution of the GMTED2010 data set is 1 km, the 3" gap-filled Shuttle Radar 139 Topography Mission (SRTM) data set (Farr et al., 2007) is used for domain 2. As this resolution is still too coarse 140 141 for our targeted resolution of 50 m, the Digital Elevation model Europe (EU-DEM; European Union, 2017), 142 available at a resolution of 25 m, is used for the innermost domain.

In our set-up, we use 100 vertical levels for all domains using the traditional terrain following coordinate system in WRF; 20 of the levels are distributed in the lowest 1100 m above ground level (AGL). All domains apply the Noah-MP land surface model (Niu et al., 2011; Yang et al., 2011), the revised MM5 surface layer scheme based on Monin-Obukhov similarity theory (Jiménez et al., 2012), the Thompson 2-moment cloud microphysics scheme (Thompson et al., 2008) and the Rapid Radiative Transfer Model for GCMs (RRTMG; Iacono et al., 2008) for parametrizing longwave and shortwave radiation. Due to the coarser resolution of the outermost domain, we applied the Yonsei University (YSU; Hong et al., 2006) planetary boundary layer (PBL) parametrization in D01





only. As suggested by the WRF user guide, we applied the sub-grid turbulent stress option for momentum(Kosovic, 1997) in domains two and three. The complete namelist settings are provided in the supplement.

152 As the finest resolution applied for the AQFS is 50 m, the more sophisticated Building Effect Parameterization

153 (BEP; Martilli et al., 2002) is not applied. as this scheme does not work with our selection of parametrizations.

154 Instead, the single layer urban canopy model (UCM) (Kusaka and Kimura, 2004) is selected to improve the

representation of the urban canopy layer and the surface fluxes. The parameters needed by the UCM are read in

156 from the lookup table URBPARAM.TBL which was adjusted for the Stuttgart area following Fallmann (2014).

157 Atmospheric chemistry is parametrized by the Regional Acid Deposition Model 2nd generation (RADM2) model (Stockwell et al., 1990). RADM2 features 63 chemical species including photolysis and more than reactions. 158 159 Aerosols are represented by the Modal Aerosol Dynamics Model for Europe (MADE) and Secondary Organic Aerosol Model (SORGAM) scheme (Ackermann et al., 1998; Schell et al., 2001) considering size distributions, 160 161 nucleation, coagulation, and condensational growth. The combination of RADM2_MADE-SORGAM is a 162 computationally efficient approach and is widely used for simulations over Europe (Forkel et al., 2015; Mar et al., 2016). To further enhance vertical mixing of CO to higher altitudes during nighttime over urban grid cells, the if-163 164 statements in the dry deposition driver of WRF-Chem at lines 690 and 707 have been deleted according as shown 165 in the supplement of Kuik et al. (2018).

166 Due to the complexity of the chemistry model in combination with the very high horizontal resolution and the 167 calm meteorological conditions, the adaptive model time step option was chosen instead of a fixed time step. 168 Model output is available in 5 min intervals for the innermost model domain.

Our single day case study on the turbulence permitting (TP) scale is designed to serve as a test bed to set up an air quality forecasting system prototype for the Stuttgart metropolitan area. For process studies, the model chain itself can be applied to other areas over the globe as long as 1) detailed land cover and soil texture data are available, 2) high-resolution emission data not only from traffic are available. The new model system can be even applied in a forecast and warning mode, if near real time emission data exist. As the computational demands of applying WRF-Chem on the TP scale are very high, access to an HPC system is a prerequisite.

175 2.2. Model initialization

The meteorological initial and boundary conditions were provided by the operational ECMWF integrated
forecasting system (IFS) analysis on model levels. The IFS is a global model with 9 km horizontal resolution and
applies a sophisticated four-dimensional variational (4DVAR) data assimilation system (Bonavita et al., 2016).
The data have been retrieved from the ECMWF Meteorological Archival and Retrieval System (MARS) and were
interpolated to a resolution of 0.05°.

181 The initialization and provision of the boundary conditions of the chemistry of the model is done with data from 182 the Whole Atmosphere Community Climate Model (WACCM; Marsh et al., 2013) using the Model for Ozone and 183 Related Chemical Tracers (MOZART) conversion tool MOZBC (Pfister et al., 2011). As the resolution of 184 WACCM is very coarse, the input data was extended by the ECMWF Copernicus Atmosphere Monitoring Service 185 (CAMS) reanalysis data set on 60 model levels and 40 km horizontal resolution (Inness et al., 2019).

186 2.3. Emission data





- The emission data set used in this study is a combination of three products. Global input data sets containing
 coarse resolution emissions from different sources are obtained from the BRAMS numerical modeling system
- 189 (Freitas et al., 2017). The PREP-CHEM-SRC tool (Freitas et al., 2011) is then applied as pre-processor to convert
- 190 these emissions to the appropriate WRF units and interpolate the data onto the WRF model grid.
- 191 As global emission data sets have a very coarse resolution in space and time, higher resolution emission data for
- Europe from the Copernicus Atmosphere Monitoring Service (CAMS; Copernicus) CAMS-REG-AP product
 became available (Granier et al., 2019). Its resolution is approx. 7x7 km and it is based on total annual emissions
 from 2016. This product provides emissions of PM₁₀, PM_{2.5}, SO₂, CO, NO_x, and CH₄ and contains sources from
 different sectors, separated into ten different categories following the Gridded Nomenclature For Reporting
 (GNFR; Granier et al., 2019).
- 197 The third emission data set (BW-EMISS) deployed in our study was obtained from the Baden-Württemberg State 198 Institute for the Environment (LUBW). This data set contains annual mean emissions from different sectors 199 following the GNFR classification and is currently available only until 2014 and has a horizontal resolution of 500 m. Unfortunately, more recent quality-controlled data sets were not available when our study was performed. It is 201 expected that annual emissions for 2018 will become available by mid of 2021.
- As CAMS-REG-AP and BW-EMISS only contain annual sums or annual mean values, a temporal decomposition
 was applied for both data sets following Denier van der Gon et al. (2011). Depending on the GNFR code, the data
 are first projected onto the corresponding month, followed by the corresponding day of the week and the hour of
 the day. A similar approach was performed e.g. in Resler et al. (2020, under review) for the city of Prague. After
 finishing the decomposition, the data are converted to the corresponding units and interpolated onto the WRF
 model grid using the Earth System Model Framework (ESMF; Valcke et al., 2012) interpolation utilities.
- Figure 3 shows an example of the NO₂ emissions derived from the CAMS-REG-AP product (left) and the emission
 data derived from the LUBW data set (right) on January 21, 2019 at 07 UTC.
- 210 Due to its much higher horizontal resolution, the BW-EMISS data set (Fig. 3b) shows much more detailed211 structures for the NO₂ emissions which are mainly caused by road traffic.
- 212 In addition, the following adjustments have been performed: 1) NO_x emissions from forest grid cells have been
- 213 reduced by 90 %, 2) Road traffic NOx emissions were transformed into 90 % NO and 10 % NO2 emissions
- following Kuik et al. (2018) 3) All emissions from Stuttgart airport were reduced by 90 % during the nighttime
- 215 flight ban between 00 UTC and 04 UTC as well as after 21 UTC.
- 216 The WRF-Chem model only ingests one emission data set per species, hence emissions from the different GNFR 217 categories have been accumulated to a single emission data set before performing the simulation. Figure 4 218 summarizes all necessary steps and the complete data and workflow of the AQFS prototype.

219 2.4. Observations

- 220 We used data from three meteorological stations (Stuttgart-Schnarrenberg (48.8281°N 9.2°E, elevation 314 m),
- 221 Stuttgart Airport (48.6883°N 9.2235°E, elevation 375 m), and Institute of Physics and Meteorology (IPM) at the
- 222 University of Hohenheim (48.716°N 9.213°E, elevation 407 m) to validate the simulated 2-m temperatures; data





are available every 10 minutes. The locations are indicated by the black dots Fig. 1b. In addition, the radiosondedata from Stuttgart-Schnarrenberg were used.

225 As the incorporated emissions are from 2014 and are based on annual values, it would be rather misleading to 226 compare the observed pollutant concentrations directly with the model output. For instance, the actual traffic, the 227 sequence of traffic lights and traffic congestions of this particular day cannot be realistically represented. In 228 addition, all diagnosed or prognostic chemical quantities are only available on model levels (with the lowest model half level being at ~15 m above ground) so that simulated concentrations need to be interpolated to the 229 230 measurement heights at 2.5-3.5 m AGL. This extrapolation may cause even more uncertainty. Therefore we 231 decided not to use pollution measurements for model evaluation but we studied the results based on process 232 understanding and plausibility arguments.

233 3. Case study description

For our study, we selected 21 January 2019. This day was characterized as "fine dust alarm" situation (Stuttgart
Municipality and German Meteorological Service (DWD), 2019) which is defined by a combination of the
following criteria:

- Expected daily maximum PM₁₀ concentration at Stuttgart Neckartor (NT in Fig. 1b) is higher than 30 μg
 m⁻³
- 239 2. No rain on the following day
- 240 3. 10-m wind speed less than 3 m s⁻¹ from south to northwest directions (180-330 $^{\circ}$)
- 241 4. Nocturnal atmospheric inversion
- 5. Mixing layer depth less than 500 m during the day
- 243 6. Daily average 10-m wind speed less than 3 m s⁻¹ from all directions

A sufficient criterion is a higher PM_{10} concentration following (1). If (1) is not fulfilled, then (2) and (3) together with either (4) and/or (5) must be fulfilled. If only (4) or (5) is fulfilled, then (6) must be considered. For our case study, the criteria 1-5 were fulfilled.

Figure 5 shows the observed PM_{10} and NO_2 concentrations at several stations in our model domain. From Fig. 5a the high NO_2 concentrations at Neckartor and Hohenheimer Strasse occurring after sunrise can be clearly identified. While these measurements are taken next to main roads, the other stations show considerably lower NO_2 concentrations throughout the day. The PM_{10} concentrations (Fig. 5b) show extremely high values at Neckartor exceeding $100 \,\mu g \, m^{-3}$ around noon time and the evening rush hour which clearly meets the main criteria of the "fine dust alarm situation". The other stations, which are not directly taken near main roads with heavy traffic show considerably lower PM_{10} concentrations around $40 \,\mu g \, m^{-3}$.

This day was a typical winter weather situation. Central Europe was located at the east flank of a blocking high pressure system located over the East Atlantic together with moderate to low horizontal geopotential gradients and resulting weak winds at 500 hPa in southwestern Germany (Fig. 6a).

Near surface temperatures are below freezing level, between 1000 and 850 hPa very light easterly winds characterize the flow, and a dry layer is present around 925 hPa (Fig. 6b). Above 850 hPa, the wind direction rapidly changes to westerly directions, but the wind speeds remain below 5 m s⁻¹ (see Fig. 7a).





- The inversion between the two air masses inhibits vertical mixing leading to higher concentrations of aerosols inthe lowest few hundred meters above ground (AGL) and preventing air mass exchange aloft. This inversion is
- 262 further enhanced by the special orography of Stuttgart city (see later Fig. 15).
- 263 4. Results and Discussion

264 4.1. Meteorological quantities

A Skew-T diagram of the observed and simulated temperature, dew point, and wind profiles allows to evaluate the
 stratification conditions of the model. Figure 7a shows the vertical profile of the model initial conditions at
 Stuttgart-Schnarrenberg valid at 00 UTC 21 January 2019 in comparison with the observations.

The initial conditions agree well with the sounding showing a weak temperature inversion around 900 hPa with high relative humidity values up to 650 hPa. The observed and simulated lifting condensation level is 940 hPa and the integrated water vapor (PWAT) is 8 mm. Wind speed and direction agree with the observations showing a

wind shear above 850 hPa associated with low wind speeds of less than 5 m s⁻¹.

To further evaluate the stratification conditions during the day, Figure 7b shows the observed and simulated temperature, dew point, and wind profiles at 11 UTC. The vertical structure of the observation and the simulation has an almost perfect agreement. The temperature inversion layer at 910 hPa is well captured although the simulated temperatures below the inversion are too high by about 1.5 K. The humidity profile (expressed as dewpoint profile) is also very well captured with the largest moisture content below 870 hPa. Wind speed and direction above 850 hPa agree well with the observation throughout the atmosphere. In regard of the vertical model resolution, the wind situation in the lowest 1000 m AGL is also reasonably represented.

Figure 8 exemplarily shows the simulated 2-m temperature together with 10-m wind velocities at 12 UTC (noontime) to display the complexity of the Stuttgart metropolitan area.

The 2-m temperatures show a daytime warming of downtown Stuttgart and the Neckar Valley while still temperature slightly below 0°C are present at higher elevations (blue colors in Fig. 8). The wind situation is very complex due to weak wind speeds in combination with a shallow boundary layer (see later Fig. 16) but the wind flow along the upper Neckar river (south of 48.75°) is strongly pronounced. After sunset, wind speed starts to increase and the channeling effect along the Neckar weakens (not shown).

Figure 9 shows an evaluation of the diurnal cycle of 2-m temperatures at the three measurement sites Schnarrenberg, IPM and airport. Sunrise is at 07 UTC and sunset at 16 UTC and the model data are averaged over grid cells around the measurement site. The northern station Schnarrenberg shows a lower temperature throughout the day than the other two stations, which are situated 3 km apart at a similar elevation. The temperature is about 1 K colder during the day and 0.5 K colder during the night.

At Schnarrenberg, the observed diurnal cycle is reasonably well simulated with WRF. Between 00 and 15 UTC, a warm temperature bias of 1 K is present in the simulation, which turns into a small negative bias after sunset. At IPM, the simulation shows a cold bias until 04 UTC turning into a warm bias as the strong temperature drop is not simulated until 06:30 UTC. After 09 UTC until sunset the simulated temperature agrees well with the observations while later a cold bias of around 1 K is present.





For the airport station, the model stays too warm with a positive bias of almost 2 K between 05 and 09 UTC.
During the further course of the day, the bias reduces to 1 K at noon while after sunset it turns into a negative bias
of 1 K.

299 A possible reason for the larger differences at the airport and IPM before (after) sun rise (sun set) is the occurrence 300 of low stratus or fog. At the beginning of the simulation, cloud coverage were reported by 5-7 octas (broken 301 clouds) over Schnarrenberg and the airport at approx. 500 m AGL (not shown) while after 04 UTC the low level 302 clouds started to diminish at Schnarrenberg first leading to a strong cooling until the early morning which is seen 303 as a temperature drop in the observations shown in Fig. 9. This temperature drop at Schnarrenberg and IPM is also 304 simulated but with a delay of approx. 2 h. During the evening transition and the following night, the low stratus 305 is developing again at the measurement sites with a ceiling of 500 m AGL but is not simulated and thus contributes 306 to a stronger cooling in the model. Another contributing factor to the delayed cloud dissipation could be the 307 turbulence spin-up time (Kealy et al., 2019), but this is beyond the scope of this study.

308 Although no measurements of sensible heat and ground heat fluxes are available, diurnal cycles of the fluxes at 309 the three locations IPM, Schnarrenberg, and airport were investigated. Figure 10 shows the simulated surface 310 sensible heat and ground heat flux at the three different meteorological measurement sites.

The sensible heat flux (Fig. 10a) shows a typical diurnal cycle with fluxes around zero before (after) sunrise (sunset). During the day, the model simulates typical wintertime sensible heat fluxes between 40 and 100 W/m² (e.g. Zieliński et al., 2018), which nicely shows a dependency on the different underlying land cover types. Lower sensible heat fluxes occur over the sparsely vegetated surface at the airport as compared to the cropland station IPM. As the algorithm to diagnose the 2-m temperature in NOAHMP is rather complex, no clear correlation between SH and the 2-m temperature shown in Fig. 9 can be made. The latent heat fluxes (not shown) are almost zero at Schnarrenberg and less than 10 W m⁻² at the other two locations due to cold and dry winter conditions

The simulated ground heat flux (Fig. 10b) shows an interesting behavior. Until sunrise, the simulated GRDFLX at the airport and IPM shows fluctuations around -50 W m⁻² indicating some low levels clouds in accordance with the too high simulated 2-m temperatures shown in Fig. 9. During the further course of the day, IPM and airport show a clear diurnal cycle with maximum values between 100 and 170 W m⁻² reflected in the highest surface temperatures during the day (not shown).

At Schnarrenberg, most of the time the ground heat flux is less than zero indicating a cooling of the soil, while between 12 UTC and 16 UTC small positive values are simulated. As Schnarrenberg is categorized as low intensity residential (category 31) with an urban fraction of 0.5 and the UCM is applied here, energy is mainly stored in the urban canopy layer instead of being transferred into the soil.

As this day was characterized by a shallow PBL and a temperature inversion, it is worth to investigate the PBL
evolution during the day. Figure 11 shows time-height cross sections of potential temperature at IPM (top) and
Schnarrenberg (bottom).

Both locations are characterized by a very stable shallow boundary layer until 09 UTC with a depth of less than
200 m. Between 03 and 09 UTC the temperatures at Schnarrenberg are up to 1.5 K colder near the surface (Fig. 9)
resulting in a stronger potential temperature gradient up to 400 m AGL compared to the IPM location. During the

day, the boundary layer height increases to 200--400 m above ground as indicated by the constant potential





temperature (e.g. Bauer et al., 2020) which is a typical value for European winter conditions (Seidel et al., 2012;
Wang et al., 2020). The PBL height estimates are confirmed by calculating the gradient Richardson number (Ri;

Chan, 2008) (not shown) which exceeds 0.25 at this altitude (Seidel et al., 2012; Lee and Wekker, 2016). After

sunset around 15:30 UTC the boundary layer collapses to a night-time stable boundary layer and a temperature

inversion occurs again.

339 4.2. Air quality

340 The most relevant air pollutants for air quality considerations in cities are NO_2 and PM_{10} . Sources for these are

- 341 mainly truck supply, transit, and commuter traffic through the city as well as advection from motorways south,
- west, and northwest of Stuttgart. We start with the discussion of the simulated horizontal distributions followed
 by vertical cross sections of NO₂ and PM₁₀.

344 4.2.1 Horizontal distribution

Figure 12 shows the horizontal distribution of the NO₂ concentration at the lowest model half level (~15 m AGL)
at the four timesteps 07:30 UTC, 12 UTC, 18 UTC and 23 UTC 21 January 2019.

347 At 7:30 UTC the morning traffic rush hour is visible in the NO₂ concentrations in Fig. 12a. High NO₂ 348 concentrations of more than 80 µg m⁻³ are simulated along the motorway A81 in the northwest of the domain, over 349 the airport and over downtown Stuttgart. In the Neckar Valley the concentrations exceed 120 µg m⁻³. At noon time 350 (Fig. 12b), when turbulence is fully evolved (Fig. 11), the simulated NO₂ concentrations are less than 30 μ g m⁻³ 351 on average apparently due to vertical mixing of NO_2 (see next section). In the evening (Fig. 12c) the simulated 352 NO₂ concentrations increase again showing values of more than 100 μ g m⁻³ over the airport and more than 150 μ g 353 m⁻³ in downtown Stuttgart and the Neckar Valley due to road and air traffic. The high morning concentrations 354 along the northwestern motorway are not reached since the wind speed increases and the near surface winds turn 355 towards a westerly direction. According to the emission data set converted by the temporal factors, the evening 356 traffic spreads over a longer time. During the night (Fig. 12d), NO2 accumulates in the Stuttgart basin as well as 357 the Neckar Valley due to the very low nocturnal boundary layer height of less than 200 m capped by an atmospheric 358 inversion (Fig. 11).

- Apart from NO₂, the concentration PM₁₀ is an important parameter for air quality considerations and is the decisive
 factor for proclaiming a "fine dust alarm" situation in Stuttgart (Stuttgart Municipality and German Meteorological
 Service (DWD), 2019). Note that the simulated PM₁₀ concentrations include particles less than 2.5 µm of diameter
- $\label{eq:2.5} \textbf{362} \qquad (PM_{2.5}) \text{ and that } PM_{10} \text{ is a diagnosed quantity in our model setup.}$
- $\label{eq:source} 363 \qquad \mbox{Figure 13 shows the horizontal distribution of PM_{10} for the same time steps as shown in Fig 12.}$

364 During the morning traffic (Fig. 13a), PM₁₀ accumulates in the Stuttgart basin as this is an area with heavy traffic 365 during the morning and an atmospheric inversion is present (Fig. 7). Interestingly, the high NO₂ concentrations 366 along the motorway (Fig. 12a) do not lead to very high PM₁₀ concentrations potentially due to chemical transitions 367 caused by low temperatures.

- 368 During daytime when turbulence is fully evolved, the concentration of PM_{10} decreases to less than 20 µg m⁻³ due 369 to vertical mixing and horizontal transport (see next section). After sunset (Fig. 13c) PM10 starts to accumulate 370 again in the Stuttgart basin showing concentrations between 35—40 µg m⁻³. During the night (Fig. 13d) PM₁₀ 371 accumulates over a large part of the model domain as the nocturnal boundary layer is very shallow, an inversion
- area layer is present 400 m AGL and the wind direction changes from north to west.





373 4.2.2 Vertical distribution of NO₂ and PM₁₀

In addition to the horizontal distribution of near surface NO_2 and PM_{10} , TP simulations with a fine vertical resolution also enable qualitative insights into the vertical distribution of pollutants. Figure 14 shows West-East cross sections at Neckartor (Fig. 1b) during the morning rush hour and at noon time. Neckartor is one of the heaviest traffic locations in the Stuttgart city area.

378 The NO₂ concentration during the morning rush hour shows an accumulation along the motorway (red arrow in

Fig. 14a) and in the region around Neckartor (white arrow in Fig. 14a) with concentrations exceeding $100 \ \mu g \ m^{-3}$

as the atmospheric inversion prevents exchange with the layers above (Fig. 7). The vertical extent of concentrations

381 higher than 30 μ g m⁻³ is about 200 m AGL with a strong reduction above.

382 During noon time (Fig. 14b), the simulated NO₂ concentration is much lower (less than 30 μ g m⁻³) as turbulence 383 leads to a stronger mixing throughout the boundary layer up to 400 m AGL which is in accordance with the 384 simulated potential temperature timeseries shown in Fig. 11.

 $\label{eq:states} \textbf{385} \qquad \text{Figure 15a displays the simulated PM}_{10} \text{ concentrations during the morning rush hour. Similar like for NO_2, higher}$

 $\label{eq:concentrations} 386 \qquad \text{concentrations of more than } 25\,\mu\text{g}\,\text{m}^{\text{-}3}\,\text{is simulated along the motorway and in the Stuttgart basin. During the day,}$

387 PM_{10} is vertically mixed showing a clear gradient around 800 m above sea level (ASL) (Fig. 15b) while

388 concentrations remain between 10-20 $\mu g \ m^{\text{-}3}$ within the boundary layer.

 $\label{eq:approx} \textbf{Apart from the West-East cross sections it is also worthwhile to investigate the vertical temporal evolution of NO_2$

and PM₁₀ concentrations. Therefore, Fig. 16 shows time height cross sections of NO₂ (top) and PM₁₀ (bottom) at
 Neckartor as an example for a heavy traffic area in the basin.

Well visible are the high simulated NO₂ and PM₁₀ concentrations during the morning rush hour with peak values
 of more than 120 µg m⁻³ NO₂ and more than 40 µg m⁻³ PM₁₀. The high concentrations of NO₂ and PM₁₀ are present

up to around 150-200 m AGL. During daytime, turbulence efficiently mixes the pollutants up to higher altitude
and the near surface concentrations are quickly reduced. During the evening when the very shallow boundary layer
has developed again and evening traffic commences, the particle concentrations increase, and peak values of more
than 30 µg m⁻³ are simulated below 100 m AGL.

398

399 5. Summary and conclusion

This paper describes the setup of a AQFS prototype using WRF-Chem for the Stuttgart Metropolitan area. Because
of the complex topography in this region, this simulation system requires a very high horizontal resolution down
to the turbulence permitting scale to represent all orographic and land cover features.

For the development of this prototype 21 January 2019 served as test case as this was a typical winter day with an atmospheric inversion. In addition, this day was characterized as "fine dust alarm" situation where the PM_{10} concentration at one of the heaviest traffic areas in the Stuttgart basin was expected to exceed 30 µg m⁻³. The model setup encompassed three domains down to a turbulence permitting resolution of 50 m.

407 The initial conditions were provided by the ECMWF operational analysis, the CAMS reanalysis and WACCM 408 model for background chemistry. Emission data sets from CAMS-REG-AP and high-resolution data with 500 m 409 resolution from LUBW were combined to be used in the AQFS. As current emission data sets only provide annual

410 totals or means, a temporal decomposition following TNO was applied (Denier van der Gon et al., 2011).





For this case study, we focused on the results with respect to 2-m temperature, surface fluxes and boundary layer
evolution as well as horizontal and vertical distributions of NO₂ and PM₁₀.

413 Our results revealed that despite the complex topography in Stuttgart, the model is in general capable to simulate 414 a realistic diurnal cycle of 2-m temperatures although, compared to observations, differences of up to 1 K occur. 415 Apparently the model has difficulties with the dissolution of low stratus clouds between 03 and 06 UTC which 416 was also reported in the work of Steeneveld et al. (2015) resulting in a warm 2-m temperature bias during the 417 morning. Although no measurements are available, the surface sensible heat fluxes show a clear diurnal cycle with 418 the magnitude clearly depending on the underlying land cover type. The low simulated ground heat flux and its 419 fluctuations between 00 UTC and sunrise partially confirm the fog dissolution issue but more test cases are needed 420 for a more detailed investigation. Over grid cells where the single layer UCM is active, most of the ground heat 421 flux is stored in the canopy layer thus not transferred into the soil. The high vertical resolution of 100 levels enables 422 a realistic representation of the nocturnal and daytime temperature inversion with an accompanying shallow 423 boundary layer of less than 400 m during the day.

The simulation of PM_{10} shows an exceedance of the 30 µg m⁻³ concentration threshold at the Neckartor station and also fulfills the other fine dust alarm criteria shown in section 3. Compared to the usually unevenly distributed air quality measurements, the AQFS allows further insights into the spatio-temporal pollutant distribution. The horizontal distributions of NO₂ and PM₁₀ at this particular day clearly indicate the main polluted areas along the motorways and in the Stuttgart basin. The special orography of Stuttgart with its basin favors the accumulation of NO₂ and PM₁₀ in the morning and evening while the pollutants are well mixed to around 200-400 m AGL when the boundary layer is fully evolved.

The simulation also shows that pollutants can be advected from the motorway A81 towards Stuttgart, depending
on the wind situation potentially leading to an increase of the NO₂ and PM₁₀ concentrations in the Stuttgart basin.
As can be seen from Figs. 12 and 13, the Neckar Valley can also have a large impact on the pollutant concentration
in the Stuttgart basin in case an atmospheric inversion together with prevailing easterly winds is present.

This is, to our knowledge, the first study of applying WRF-Chem on a TP resolution for an urban area. To derive
more robust conclusions with respect to air pollution, more cases studies with different weather situations during
winter and summer time are necessary. Nevertheless, our evaluation gives the following indications to further
improve the quality of such simulations:

- 439 I. Applying high spatial and temporal resolution gridded emission data from all pollution sources in near
 440 real time to avoid extrapolating annual emissions to individual days.. This will help to enhance the
 441 simulation of the diurnal cycles of chemical species.
- II. Improving the chemical background e.g. by applying higher resolution products from the CAMS
 European Air quality project (Marécal et al., 2015). This will help to have a more detailed structure of
 the chemical constituents beneficial for further downscaling simulations.
- 445 III. Using a longer spin-up period and applying a larger TP model domain to further improve the spin-up of446 turbulence in the model
- 447 IV. Considering vertical distribution of surface emissions (e.g. Bieser et al., 2011; Guevara et al., 2020)
- 448 V. Considerably increase the number of pollutant measurements to allow more robust conclusions





The AQFS has a great potential for urban planning applications. For example, land cover could be changed from urban low density to urban high density to investigate the impact of urban re-densification e.g. on temperature and air quality. Although no BEP can be applied on the TP resolution with our combination of parameterizations, changes of the parameters required for the single layer UCM offer the opportunity to perform sensitivity analysis with respect to different building heights, urban greening effects (Fallmann et al., 2016), or anthropogenic heat (Karlický et al., 2020).

Although air quality modeling on the TP scale is a very challenging and computationally expensive task, we are
convinced that the AQFS will have a great potential to further improve process understanding and will certainly
help politicians to make decisions on a more scientifically valid basis.

458 Code and data availability

459 WRF-Chem code 4.0.3 The version can be downloaded from https://github.com/wrfmodel/WRF/archive/v4.0.3.tar.gz. ECMWF analysis data can be obtained from https://apps.ecmwf.int/archive-460 461 catalogue/?type=an&class=od&stream=oper&expver=1 (last access: 26 August 2020). The user's affiliation needs 462 to belong to an ECMWF member state to benefit from these data sets. Due to restrictions on the input data sets for 463 this simulation, the data can only be made available upon special request from the corresponding author.

464 Author Contributions

TS prepared all emission data, set up the model and performed the simulation supported by HSB. HSB reclassified
the CORINE land use data set. KWS and TB conceived the idea and coordinated the project with VW. TS prepared
all figures and wrote the manuscript with input from all authors. All authors equally contributed to the scientific
discussion and helped to shape the research.

469 Competing interests

470 The authors declare that they have no conflict of interest.

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484 References





485 486	Ackermann, I. J., Hass, H., Memmesheimer, M., Ebel, A., Binkowski, F. S., and Shankar, U.: Modal aerosol dynamics model for Europe, Atmospheric Environment, 32, 2981–2999,
487	https://doi.org/10.1016/S1352-2310(98)00006-5, 1998.
488	Baldauf, M., Seifert, A., Förstner, J., Majewski, D., Raschendorfer, M., and Reinhardt, T.:
489	Operational Convective-Scale Numerical Weather Prediction with the COSMO Model:
490	Description and Sensitivities Mon Wea Rev. 139 3887–3905
491	https://doi.org/10.1175/MWR-D-10-05013.1.2011
492	Barker D. Huang X-Y. Liu, Z. Auligné T. Zhang X. Rugg S. Aijaji R. Bourgeois A. Bray
493	J., Chen, Y., Demirtas, M., Guo, YR., Henderson, T., Huang, W., Lin, HC., Michalakes, J.,
494	Rizvi, S., and Zhang, X.: The Weather Research and Forecasting Model's Community
495	Variational/Ensemble Data Assimilation System: WRFDA, Bull. Amer. Meteor. Soc., 93,
496	831–843, https://doi.org/10.1175/BAMS-D-11-00167.1, 2012.
497	Bauer, HS., Muppa, S. K., Wulfmeyer, V., Behrendt, A., Warrach-Sagi, K., and Späth, F.:
498	Multi-nested WRF simulations for studying planetary boundary layer processes on the
499	turbulence-permitting scale in a realistic mesoscale environment, Tellus A: Dynamic
500	Meteorology and Oceanography, 72, 1–28,
501	https://doi.org/10.1080/16000870.2020.1761740, 2020.
502	Bieser, J., Aulinger, A., Matthias, V., Quante, M., and van der Denier Gon, H. A. C.: Vertical
503	emission profiles for Europe based on plume rise calculations, Environmental pollution
504	(Barking, Essex 1987), 159, 2935–2946, https://doi.org/10.1016/j.envpol.2011.04.030,
505	2011.
506	Bonavita, M., Hólm, E., Isaksen, L., and Fisher, M.: The evolution of the ECMWF hybrid data
507	assimilation system, Q.J.R. Meteorol. Soc., 142, 287–303,
508	https://doi.org/10.1002/qj.2652, 2016.
509	Brasseur, G. P., Hauglustaine, D. A., Walters, S., Rasch, P. J., Müller, JF., Granier, C., and Tie,
510	X. X.: MOZART, a global chemical transport model for ozone and related chemical tracers:
511	1. Model description, J. Geophys. Res., 103, 28265–28289,
512	https://doi.org/10.1029/98JD02397, 1998.
513	Butcher, G., Barnes, C., and Owen, L.: Landsat: The cornerstone of global land imaging, GIM
514	International, January/February 2019, 31–35, available at:
515	http://pubs.er.usgs.gov/publication/70202363, 2019.
516	Chan, P. W.: Determination of Richardson number profile from remote sensing data and its
517	aviation application, IOP Conf. Ser.: Earth Environ. Sci., 1, 12043,
518	https://doi.org/10.1088/1755-1315/1/1/012043, 2008.
519	Chen, D., Liu, Z., Ban, J., Zhao, P., and Chen, M.: Retrospective analysis of 2015–2017
520	wintertime $PM_{2.5}$ in China: response to emission regulations and the role of meteorology,
521	Atmos. Chem. Phys., 19, 7409–7427, https://doi.org/10.5194/acp-19-7409-2019, 2019.
522	Copernicus: Copernicus official website, https://atmosphere.copernicus.eu/, last access: 21
523	July 2020.
524	Coppola, E., Sobolowski, S., Pichelli, E., Raffaele, F., Ahrens, B., Anders, I., Ban, N., Bastin, S.,
525	Belda, M., Belusic, D., Caldas-Alvarez, A., Cardoso, R. M., Davolio, S., Dobler, A.,
526	Fernandez, J., Fita, L., Fumiere, Q., Giorgi, F., Goergen, K., Güttler, I., Halenka, T.,
527	Heinzeller, D., Hodnebrog, Ø., Jacob, D., Kartsios, S., Katragkou, E., Kendon, E., Khodayar,
528	S., Kunstmann, H., Knist, S., Lavín-Gullón, A., Lind, P., Lorenz, T., Maraun, D., Marelle, L.,





529	van Meijgaard, E., Milovac, J., Myhre, G., Panitz, HJ., Piazza, M., Raffa, M., Raub, T.,
530	Rockel, B., Schär, C., Sieck, K., Soares, P. M. M., Somot, S., Srnec, L., Stocchi, P., Tölle, M.
531	H., Truhetz, H., Vautard, R., Vries, H. de, and Warrach-Sagi, K.: A first-of-its-kind multi-
532	model convection permitting ensemble for investigating convective phenomena over
533	Europe and the Mediterranean, Clim Dyn, 55, 3–34, https://doi.org/10.1007/s00382-018-
534	4521-8, 2020.
535	Corsmeier, U., Kalthoff, N., Barthlott, C., Aoshima, F., Behrendt, A., Di Girolamo, P.,
536	Dorninger, M., Handwerker, J., Kottmeier, C., Mahlke, H., Mobbs, S. D., Norton, E. G.,
537	Wickert, J., and Wulfmeyer, V.: Processes driving deep convection over complex terrain:
538	a multi-scale analysis of observations from COPS IOP 9c, Q.J.R. Meteorol. Soc., 137, 137–
539	155, https://doi.org/10.1002/qj.754, 2011.
540	Danielson, J. J. and Gesch, D. B.: Global multi-resolution terrain elevation data 2010
541	(GMTED2010), Open-File report, 2011.
542	Denier van der Gon, H., Hendriks, C., Kuenen, J., Segers, A., and Visschedijk, A.: Description
543	of current temporal emission patterns and sensitivity of predicted AQ for temporal
544	emission patterns: EU FP7 MACC deliverable report D_D-EMIS_1.3, TNO report, 2011.
545	European Union: Copernicus Land Monitoring Service 2017, European Environment Agency,
546	2017.
547	European Union: Copernicus Land Monitoring Service 2012, European Environment Agency,
548	2012.
549	Fallmann, J.: Numerical simulations to assess the effect of urban heat island mitigation
550	strategies on regional air quality, PhD Thesis, Universität zu Köln, Cologne, 2014.
551	Fallmann, J., Forkel, R., and Emeis, S.: Secondary effects of urban heat island mitigation
552	measures on air quality, Atmospheric Environment, 125, 199–211,
553	https://doi.org/10.1016/j.atmosenv.2015.10.094, 2016.
554	Fallmann, J., Emeis, S., and Suppan, P.: Mitigation of urban heat stress – a modelling case
555	study for the area of Stuttgart, DIE ERDE – Journal of the Geographical Society of Berlin,
556	144, 202–216, https://doi.org/10.12854/erde-144-15, 2014.
557	Farr, T. G., Rosen, P. A., Caro, E., Crippen, R., Duren, R., Hensley, S., Kobrick, M., Paller, M.,
558	Rodriguez, E., Roth, L., Seal, D., Shaffer, S., Shimada, J., Umland, J., Werner, M., Oskin,
559	M., Burbank, D., and Alsdorf, D.: The Shuttle Radar Topography Mission, Rev. Geophys.,
560	45, https://doi.org/10.1029/2005RG000183, 2007.
561	Flemming, J., Huijnen, V., Arteta, J., Bechtold, P., Beljaars, A., Blechschmidt, AM.,
562	Diamantakis, M., Engelen, R. J., Gaudel, A., Inness, A., Jones, L., Josse, B., Katragkou, E.,
563	Marecal, V., Peuch, VH., Richter, A., Schultz, M. G., Stein, O., and Tsikerdekis, A.:
564	Tropospheric chemistry in the Integrated Forecasting System of ECMWF, Geosci. Model
565	Dev., 8, 975–1003, https://doi.org/10.5194/gmd-8-975-2015, 2015.
566	Forkel, R., Balzarini, A., Baró, R., Bianconi, R., Curci, G., Jiménez-Guerrero, P., Hirtl, M.,
567	Honzak, L., Lorenz, C., Im, U., Pérez, J. L., Pirovano, G., San José, R., Tuccella, P., Werhahn,
568	J., and Žabkar, R.: Analysis of the WRF-Chem contributions to AQMEII phase2 with
569	respect to aerosol radiative feedbacks on meteorology and pollutant distributions,
570	Atmospheric Environment, 115, 630–645,
571	https://doi.org/10.1016/j.atmosenv.2014.10.056, 2015.





572 Freitas, S. R., Longo, K. M., Alonso, M. F., Pirre, M., Marecal, V., Grell, G., Stockler, R., Mello, 573 R. F., and Sánchez Gácita, M.: PREP-CHEM-SRC – 1.0: a preprocessor of trace gas and 574 aerosol emission fields for regional and global atmospheric chemistry models, Geosci. 575 Model Dev., 4, 419–433, https://doi.org/10.5194/gmd-4-419-2011, 2011. 576 Freitas, S. R., Panetta, J., Longo, K. M., Rodrigues, L. F., Moreira, D. S., Rosário, N. E., Silva 577 Dias, P. L., Silva Dias, M. A. F., Souza, E. P., Freitas, E. D., Longo, M., Frassoni, A., Fazenda, 578 A. L., Santos e Silva, C. M., Pavani, C. A. B., Eiras, D., França, D. A., Massaru, D., Silva, F. B., 579 Santos, F. C., Pereira, G., Camponogara, G., Ferrada, G. A., Campos Velho, H. F., Menezes, 580 I., Freire, J. L., Alonso, M. F., Gácita, M. S., Zarzur, M., Fonseca, R. M., Lima, R. S., Siqueira, R. A., Braz, R., Tomita, S., Oliveira, V., and Martins, L. D.: The Brazilian developments on 581 582 the Regional Atmospheric Modeling System (BRAMS 5.2): an integrated environmental model tuned for tropical areas, Geosci. Model Dev., 10, 189-222, 583 584 https://doi.org/10.5194/gmd-10-189-2017, 2017. 585 García-Díez, M., Lauwaet, D., Hooyberghs, H., Ballester, J., Ridder, K. de, and Rodó, X.: 586 Advantages of using a fast urban boundary layer model as compared to a full mesoscale model to simulate the urban heat island of Barcelona, Geosci. Model Dev., 9, 4439–4450, 587 https://doi.org/10.5194/gmd-9-4439-2016, 2016. 588 589 Giorgi, F., Coppola, E., Solmon, F., Mariotti, L., Sylla, M. B., Bi, X., Elguindi, N., Diro, G. T., 590 Nair, V., Giuliani, G., Turuncoglu, U. U., Cozzini, S., Güttler, I., O'Brien, T. A., Tawfik, A. B., Shalaby, A., Zakey, A. S., Steiner, A. L., Stordal, F., Sloan, L. C., and Brankovic, C.: RegCM4: 591 model description and preliminary tests over multiple CORDEX domains, Clim. Res., 52, 592 593 7-29, https://doi.org/10.3354/cr01018, 2012. Granier, C., Darras, S., Denier van der Gon, H., Doubalova, J., Elguindi, N., Galle, B., Gauss, 594 M., Guevara, M., Jalkanen, J.-P., Kuenen, J., Liousse, C., Quack, B., Simpson, D., and 595 596 Sindelarova, K.: The Copernicus Atmosphere Monitoring Service global and regional emissions (April 2019 version), 2019. 597 Grell, G. A., Peckham, S. E., Schmitz, R., McKeen, S. A., Frost, G., Skamarock, W. C., and Eder, 598 599 B.: Fully coupled "online" chemistry within the WRF model, Atmospheric Environment, 39, 6957–6975, https://doi.org/10.1016/j.atmosenv.2005.04.027, 2005. 600 601 Guevara, M., Jorba, O., Tena, C., van der Denier Gon, H., Kuenen, J., Elguindi-Solmon, N., 602 Darras, S., Granier, C., and Pérez García-Pando, C.: CAMS-TEMPO: global and European emission temporal profile maps for atmospheric chemistry modelling, 2020. 603 Heinze, R., Moseley, C., Böske, L. N., Muppa, S. K., Maurer, V., Raasch, S., and Stevens, B.: 604 Evaluation of large-eddy simulations forced with mesoscale model output for a multi-605 606 week period during a measurement campaign, Atmos. Chem. Phys., 17, 7083-7109, 607 https://doi.org/10.5194/acp-17-7083-2017, 2017a. 608 Heinze, R., Dipankar, A., Henken, C. C., Moseley, C., Sourdeval, O., Trömel, S., Xie, X., Adamidis, P., Ament, F., Baars, H., Barthlott, C., Behrendt, A., Blahak, U., Bley, S., Brdar, 609 610 S., Brueck, M., Crewell, S., Deneke, H., Di Girolamo, P., Evaristo, R., Fischer, J., Frank, C., 611 Friederichs, P., Göcke, T., Gorges, K., Hande, L., Hanke, M., Hansen, A., Hege, H.-C., 612 Hoose, C., Jahns, T., Kalthoff, N., Klocke, D., Kneifel, S., Knippertz, P., Kuhn, A., van Laar, 613 T., Macke, A., Maurer, V., Mayer, B., Meyer, C. I., Muppa, S. K., Neggers, R. A. J., Orlandi, E., Pantillon, F., Pospichal, B., Röber, N., Scheck, L., Seifert, A., Seifert, P., Senf, F., Siligam, 614 615 P., Simmer, C., Steinke, S., Stevens, B., Wapler, K., Weniger, M., Wulfmeyer, V., Zängl, G.,





616	Zhang, D., and Quaas, J.: Large-eddy simulations over Germany using ICON: a
617	comprehensive evaluation, Q.J.R. Meteorol. Soc., 143, 69–100,
618	https://doi.org/10.1002/qj.2947, 2017b.
619	Hengl, T., Heuvelink, G. B. M., Kempen, B., Leenaars, J. G. B., Walsh, M. G., Shepherd, K. D.,
620	Sila, A., MacMillan, R. A., Mendes de Jesus, J., Tamene, L., and Tondoh, J. E.: Mapping Soil
621	Properties of Africa at 250 m Resolution: Random Forests Significantly Improve Current
622	Predictions, PloS one, 10, e0125814, https://doi.org/10.1371/journal.pone.0125814,
623	2015.
624	Hengl, T., Jesus, J. M. de, MacMillan, R. A., Batjes, N. H., Heuvelink, G. B. M., Ribeiro, E.,
625	Samuel-Rosa, A., Kempen, B., Leenaars, J. G. B., Walsh, M. G., and Gonzalez, M. R.:
626	SoilGrids1kmglobal soil information based on automated mapping, PloS one, 9,
627	e105992, https://doi.org/10.1371/journal.pone.0105992, 2014.
628	Hong, SY., Noh, Y., and Dudhia, J.: A New Vertical Diffusion Package with an Explicit
629	Treatment of Entrainment Processes, Mon. Wea. Rev., 134, 2318–2341,
630	https://doi.org/10.1175/MWR3199.1, 2006.
631	Horowitz, L. W., Walters, S., Mauzerall, D. L., Emmons, L. K., Rasch, P. J., Granier, C., Tie, X.,
632	Lamarque, JF., Schultz, M. G., Tyndall, G. S., Orlando, J. J., and Brasseur, G. P.: A global
633	simulation of tropospheric ozone and related tracers: Description and evaluation of
634	MOZART, version 2, J. Geophys. Res., 108, n/a-n/a,
635	https://doi.org/10.1029/2002JD002853, 2003.
636	Huszar, P., Karlický, J., Ďoubalová, J., Šindelářová, K., Nováková, T., Belda, M., Halenka, T.,
637	Žák, M., and Pišoft, P.: Urban canopy meteorological forcing and its impact on ozone and
638	PM _{2.5} : role of vertical turbulent transport, Atmos. Chem. Phys., 20, 1977–2016,
639	https://doi.org/10.5194/acp-20-1977-2020, 2020.
640	Iacono, M. J., Delamere, J. S., Mlawer, E. J., Shephard, M. W., Clough, S. A., and Collins, W.
641	D.: Radiative forcing by long-lived greenhouse gases: Calculations with the AER radiative
642	transfer models, J. Geophys. Res., 113, https://doi.org/10.1029/2008JD009944, 2008.
643	Inness, A., Ades, M., Agustí-Panareda, A., Barré, J., Benedictow, A., Blechschmidt, AM.,
644	Dominguez, J. J., Engelen, R., Eskes, H., Flemming, J., Huijnen, V., Jones, L., Kipling, Z.,
645	Massart, S., Parrington, M., Peuch, VH., Razinger, M., Remy, S., Schulz, M., and Suttie,
646	M.: The CAMS reanalysis of atmospheric composition, Atmos. Chem. Phys., 19, 3515-
647	3556, https://doi.org/10.5194/acp-19-3515-2019, 2019.
648	Jiménez, P. A., Dudhia, J., González-Rouco, J. F., Navarro, J., Montávez, J. P., and García-
649	Bustamante, E.: A Revised Scheme for the WRF Surface Layer Formulation, Mon. Wea.
650	Rev., 140, 898–918, https://doi.org/10.1175/MWR-D-11-00056.1, 2012.
651	Jin, L., Li, Z., He, Q., Miao, Q., Zhang, H., and Yang, X.: Observation and simulation of near-
652	surface wind and its variation with topography in Urumqi, West China, J Meteorol Res,
653	30, 961–982, https://doi.org/10.1007/s13351-016-6012-3, 2016.
654	Karlický, J., Huszár, P., Nováková, T., Belda, M., Švábik, F., Ďoubalová, J., and Halenka, T.: The
655	`urban meteorology island': a multi-model ensemble analysis, 2020.
656	Kawabata, T., Schwitalla, T., Adachi, A., Bauer, HS., Wulfmeyer, V., Nagumo, N., and
657	Yamauchi, H.: Observational operators for dual polarimetric radars in variational data
658	assimilation systems (PolRad VAR v1.0), Geosci. Model Dev., 11, 2493–2501,
659	https://doi.org/10.5194/gmd-11-2493-2018, 2018.





~~~	Kashi J. C. Efstathiau, C. A. and Daars, D. J. The Oreat of Daash ad David and Javan
660	Kealy, J. C., Erstatniou, G. A., and Beare, R. J.: The Onset of Resolved Boundary-Layer
661	Turbulence at Grey-zone Resolutions, Boundary-Layer Meteorol, 171, 31–52,
662	https://doi.org/10.100//s10546-018-0420-0, 2019.
663	Knan, B., Banzhar, S., Chan, E. C., Forkel, R., Kanani-Sunring, F., Ketelsen, K., Kurppa, M.,
664	Maronga, B., Mauder, M., Raasch, S., Russo, E., Schaap, M., and Sunring, M.:
665	Development of an atmospheric chemistry model coupled to the PALM model system
666	6.0: Implementation and first applications, Geosci. Model Dev.,
667	https://doi.org/10.5194/gmd-2020-286, 2020, under review.
668	Kosovic, B.: Subgrid-scale modelling for the large-eddy simulation of high-Reynolds-number
669	boundary layers, J. Fluid Mech., 336, 151–182,
670	https://doi.org/10.1017/S0022112096004697, 1997.
671	Kuenen, J. J. P., Visschedijk, A. J. H., Jozwicka, M., and van der Denier Gon, H. A. C.: TNO-
672	MACC_II emission inventory; a multi-year (2003–2009) consistent high-resolution
673	European emission inventory for air quality modelling, Atmos. Chem. Phys., 14, 10963–
674	10976, https://doi.org/10.5194/acp-14-10963-2014, 2014.
675	Kuik, F., Kerschbaumer, A., Lauer, A., Lupascu, A., Schneidemesser, E. von, and Butler, T. M.:
676	Top–down quantification of NO x emissions from traffic in an urban area using a high-
677	resolution regional atmospheric chemistry model, Atmos. Chem. Phys., 18, 8203–8225,
678	https://doi.org/10.5194/acp-18-8203-2018, 2018.
679	Kuik, F., Lauer, A., Churkina, G., van der Denier Gon, H. A. C., Fenner, D., Mar, K. A., and
680	Butler, T. M.: Air quality modelling in the Berlin–Brandenburg region using WRF-Chem
681	v3.7.1: sensitivity to resolution of model grid and input data, Geosci. Model Dev., 9,
682	4339–4363, https://doi.org/10.5194/gmd-9-4339-2016, 2016.
683	Kusaka, H. and Kimura, F.: Coupling a Single-Layer Urban Canopy Model with a Simple
684	Atmospheric Model: Impact on Urban Heat Island Simulation for an Idealized Case, JMSJ,
685	82, 67–80, https://doi.org/10.2151/jmsj.82.67, 2004.
686	Lee, T. R. and Wekker, S. F. J. de: Estimating Daytime Planetary Boundary Layer Heights over
687	a Valley from Rawinsonde Observations at a Nearby Airport: An Application to the Page
688	Valley in Virginia, United States, J. Appl. Meteor. Climatol., 55, 791–809,
689	https://doi.org/10.1175/JAMC-D-15-0300.1, 2016.
690	Li, Z., Zhou, Y., Wan, B., Chung, H., Huang, B., and Liu, B.: Model evaluation of high-
691	resolution urban climate simulations: using the WRF/Noah LSM/SLUCM model (Version
692	3.7.1) as a case study, Geosci. Model Dev., 12, 4571–4584, https://doi.org/10.5194/gmd-
693	12-4571-2019, 2019.
694	Mailler, S., Menut, L., Khvorostyanov, D., Valari, M., Couvidat, F., Siour, G., Turquety, S.,
695	Briant, R., Tuccella, P., Bessagnet, B., Colette, A., Létinois, L., Markakis, K., and Meleux, F.:
696	CHIMERE-2017: from urban to hemispheric chemistry-transport modeling, Geosci. Model
697	Dev., 10, 2397–2423, https://doi.org/10.5194/gmd-10-2397-2017, 2017.
698	Manders, A. M. M., Builtjes, P. J. H., Curier, L., van der Denier Gon, H. A. C., Hendriks, C.,
699	Jonkers, S., Kranenburg, R., Kuenen, J. J. P., Segers, A. J., Timmermans, R. M. A.,
700	Visschedijk, A. J. H., Wichink Kruit, R. J., van Pul, W. A. J., Sauter, F. J., van der Swaluw, E.,
701	Swart, D. P. J., Douros, J., Eskes, H., van Meijgaard, E., van Ulft, B., van Velthoven, P.,
702	Banzhaf, S., Mues, A. C., Stern, R., Fu, G., Lu, S., Heemink, A., van Velzen, N., and Schaap,





703	M.: Curriculum vitae of the LOTOS–EUROS (v2.0) chemistry transport model, Geosci.
704	Model Dev., 10, 4145–4173, https://doi.org/10.5194/gmd-10-4145-2017, 2017.
705	Mar, K. A., Ojha, N., Pozzer, A., and Butler, T. M.: Ozone air quality simulations with WRF-
706	Chem (v3.5.1) over Europe: model evaluation and chemical mechanism comparison,
707	Geosci. Model Dev., 9, 3699–3728, https://doi.org/10.5194/gmd-9-3699-2016, 2016.
708	Marécal, V., Peuch, VH., Andersson, C., Andersson, S., Arteta, J., Beekmann, M.,
709	Benedictow, A., Bergström, R., Bessagnet, B., Cansado, A., Chéroux, F., Colette, A.,
710	Coman, A., Curier, R. L., van der Denier Gon, H. A. C., Drouin, A., Elbern, H., Emili, E.,
711	Engelen, R. J., Eskes, H. J., Foret, G., Friese, E., Gauss, M., Giannaros, C., Guth, J., Joly, M.,
712	Jaumouillé, E., Josse, B., Kadygrov, N., Kaiser, J. W., Krajsek, K., Kuenen, J., Kumar, U.,
713	Liora, N., Lopez, E., Malherbe, L., Martinez, I., Melas, D., Meleux, F., Menut, L., Moinat,
714	P., Morales, T., Parmentier, J., Piacentini, A., Plu, M., Poupkou, A., Queguiner, S.,
715	Robertson, L., Rouïl, L., Schaap, M., Segers, A., Sofiev, M., Tarasson, L., Thomas, M.,
716	Timmermans, R., Valdebenito, Á., van Velthoven, P., van Versendaal, R., Vira, J., and Ung,
717	A.: A regional air quality forecasting system over Europe: the MACC-II daily ensemble
718	production, Geosci. Model Dev., 8, 2777–2813, https://doi.org/10.5194/gmd-8-2777-
719	2015, 2015.
720	Maronga, B., Gryschka, M., Heinze, R., Hoffmann, F., Kanani-Sühring, F., Keck, M., Ketelsen,
721	K., Letzel, M. O., Sühring, M., and Raasch, S.: The Parallelized Large-Eddy Simulation
722	Model (PALM) version 4.0 for atmospheric and oceanic flows: model formulation, recent
723	developments, and future perspectives, Geosci. Model Dev., 8, 2515–2551,
724	https://doi.org/10.5194/gmd-8-2515-2015, 2015.
725	Maronga, B., Banzhaf, S., Burmeister, C., Esch, T., Forkel, R., Fröhlich, D., Fuka, V., Gehrke, K.
726	F., Geletič, J., Giersch, S., Gronemeier, T., Groß, G., Heldens, W., Hellsten, A., Hoffmann,
727	F., Inagaki, A., Kadasch, E., Kanani-Sühring, F., Ketelsen, K., Khan, B. A., Knigge, C., Knoop,
728	H., Krč, P., Kurppa, M., Maamari, H., Matzarakis, A., Mauder, M., Pallasch, M., Pavlik, D.,
729	Pfafferott, J., Resler, J., Rissmann, S., Russo, E., Salim, M., Schrempf, M., Schwenkel, J.,
730	Seckmeyer, G., Schubert, S., Sühring, M., Tils, R. von, Vollmer, L., Ward, S., Witha, B.,
731	Wurps, H., Zeidler, J., and Raasch, S.: Overview of the PALM model system 6.0, Geosci.
732	Model Dev., 13, 1335–1372, https://doi.org/10.5194/gmd-13-1335-2020, 2020.
733	Maronga, B., Gross, G., Raasch, S., Banzhaf, S., Forkel, R., Heldens, W., Kanani-Sühring, F.,
734	Matzarakis, A., Mauder, M., Pavlik, D., Pfafferott, J., Schubert, S., Seckmeyer, G., Sieker,
735	H., and Winderlich, K.: Development of a new urban climate model based on the model
736	PALM – Project overview, planned work, and first achievements, metz, 28, 105–119,
737	https://doi.org/10.1127/metz/2019/0909, 2019.
738	Marsh, D. R., Mills, M. J., Kinnison, D. E., Lamarque, JF., Calvo, N., and Polvani, L. M.:
739	Climate Change from 1850 to 2005 Simulated in CESM1(WACCM), J. Climate, 26, 7372–
740	/391, https://doi.org/10.11/5/JCLI-D-12-00558.1, 2013.
/41	Martilli, A., Clappier, A., and Rotach, M. W.: An Urban Surface Exchange Parameterisation for
/42	iviesoscale Models, Boundary-Layer Meteorology, 104, 261–304,
743	nttps://doi.org/10.1023/A:1016099921195, 2002.
744	ridelinest Dest. present and future MUO Designal Office for Surger Consultance 22
745	guidelines: Past, present and future, WHO Regional Office for Europe, Copenhagen, 32
/46	pp., 2017.





747 Memmesheimer, M., Friese, E., Ebel, A., Jakobs, H. J., Feldmann, H., Kessler, C., and Piekorz, 748 G.: Long-term simulations of particulate matter in Europe on different scales using 749 sequential nesting of a regional model, IJEP, 22, 108, 750 https://doi.org/10.1504/IJEP.2004.005530, 2004. 751 Nakayama, H., Takemi, T., and Nagai, H.: Large-eddy simulation of urban boundary-layer flows by generating turbulent inflows from mesoscale meteorological simulations, 752 753 Atmosph. Sci. Lett., 13, 180–186, https://doi.org/10.1002/asl.377, 2012. 754 Niu, G.-Y., Yang, Z.-L., Mitchell, K. E., Chen, F., Ek, M. B., Barlage, M., Kumar, A., Manning, K., 755 Niyogi, D., Rosero, E., Tewari, M., and Xia, Y.: The community Noah land surface model 756 with multiparameterization options (Noah-MP): 1. Model description and evaluation with 757 local-scale measurements, J. Geophys. Res., 116, https://doi.org/10.1029/2010JD015139, 758 2011. 759 Panosetti, D., Böing, S., Schlemmer, L., and Schmidli, J.: Idealized Large-Eddy and Convection-760 Resolving Simulations of Moist Convection over Mountainous Terrain, J. Atmos. Sci., 73, 761 4021–4041, https://doi.org/10.1175/JAS-D-15-0341.1, 2016. Pfister, G. G., Parrish, D. D., Worden, H., Emmons, L. K., Edwards, D. P., Wiedinmyer, C., 762 Diskin, G. S., Huey, G., Oltmans, S. J., Thouret, V., Weinheimer, A., and Wisthaler, A.: 763 764 Characterizing summertime chemical boundary conditions for airmasses entering the US 765 West Coast, Atmos. Chem. Phys., 11, 1769–1790, https://doi.org/10.5194/acp-11-1769-766 2011, 2011. Prein, A. F., Langhans, W., Fosser, G., Ferrone, A., Ban, N., Goergen, K., Keller, M., Tölle, M., 767 768 Gutjahr, O., Feser, F., Brisson, E., Kollet, S., Schmidli, J., van Lipzig, N. P. M., and Leung, R.: 769 A review on regional convection-permitting climate modeling: Demonstrations, prospects, and challenges, Rev. Geophys., 53, 323-361, 770 771 https://doi.org/10.1002/2014RG000475, 2015. Resler, J., Eben, K., Geletič, J., Krč, P., Rosecký, M., Sühring, M., Belda, M., Fuka, V., Halenka, 772 773 T., Huszár, P., Karlický, J., Benešová, N., Ďoubalová, J., Honzáková, K., Keder, J., 774 Nápravníková, Š., and Vlček, O.: Validation of the PALM model system 6.0 in real urban environment; case study of Prague-Dejvice, Czech Republic, Geosci. Model Dev., 775 776 https://doi.org/10.5194/gmd-2020-175, 2020, under review. 777 Rieger, D., Bangert, M., Bischoff-Gauss, I., Förstner, J., Lundgren, K., Reinert, D., Schröter, J., 778 Vogel, H., Zängl, G., Ruhnke, R., and Vogel, B.: ICON-ART 1.0 - a new online-coupled 779 model system from the global to regional scale, Geosci. Model Dev., 8, 1659–1676, https://doi.org/10.5194/gmd-8-1659-2015, 2015. 780 781 Salamanca, F. and Martilli, A.: A new Building Energy Model coupled with an Urban Canopy 782 Parameterization for urban climate simulations-part II. Validation with one dimension off-line simulations, Theor Appl Climatol, 99, 345–356, https://doi.org/10.1007/s00704-783 784 009-0143-8, 2010. Schell, B., Ackermann, I. J., Hass, H., Binkowski, F. S., and Ebel, A.: Modeling the formation of 785 786 secondary organic aerosol within a comprehensive air quality model system, J. Geophys. 787 Res., 106, 28275–28293, https://doi.org/10.1029/2001JD000384, 2001. 788 Scherer, D., Antretter, F., Bender, S., Cortekar, J., Emeis, S., Fehrenbach, U., Gross, G., Halbig, G., Hasse, J., Maronga, B., Raasch, S., and Scherber, K.: Urban Climate Under Change 789 790 [UC]2 – A National Research Programme for Developing a Building-Resolving





791	Atmospheric Model for Entire City Regions, metz, 28, 95–104,
792	https://doi.org/10.1127/metz/2019/0913, 2019.
793	Seidel, D. J., Zhang, Y., Beljaars, A., Golaz, JC., Jacobson, A. R., and Medeiros, B.:
794	Climatology of the planetary boundary layer over the continental United States and
795	Europe, J. Geophys. Res., 117, n/a-n/a, https://doi.org/10.1029/2012JD018143, 2012.
796	Skamarock, W. C., Klemp, J. B., Dudhia, J., Gill, D. O., Liu, Z., Berner, J., Wang, W., Powers, J.
797	G., Duda, M. G., Barker, D. M., and Huang, XY.: A Description of the Advanced Research
798	WRF Model Version 4, 2019.
799	Steeneveld, G. J., Ronda, R. J., and Holtslag, A. A. M.: The Challenge of Forecasting the Onset
800	and Development of Radiation Fog Using Mesoscale Atmospheric Models, Boundary-
801	Layer Meteorol, 154, 265–289, https://doi.org/10.1007/s10546-014-9973-8, 2015.
802	Stockwell, W. R., Middleton, P., Chang, J. S., and Tang, X.: The second generation regional
803	acid deposition model chemical mechanism for regional air quality modeling, J. Geophys.
804	Res., 95, 16343, https://doi.org/10.1029/JD095iD10p16343, 1990.
805	Stuttgart Municipality and German Meteorological Service (DWD): Requirements for fine
806	dust situations, https://feinstaubalarm.stuttgart.de/img/mdb/item/584405/119353.pdf,
807	last access: 20 August 2020, 2019.
808	Sun, W., Liu, Z., Chen, D., Zhao, P., and Chen, M.: Development and application of the
809	WRFDA-Chem three-dimensional variational (3DVAR) system: aiming to improve air
810	quality forecasting and diagnose model deficiencies, Atmos. Chem. Phys., 20, 9311–9329,
811	https://doi.org/10.5194/acp-20-9311-2020, 2020.
812	Teixeira, J. C., Fallmann, J., Carvalho, A. C., and Rocha, A.: Surface to boundary layer coupling
813	in the urban area of Lisbon comparing different urban canopy models in WRF, Urban
814	Climate, 28, 100454, https://doi.org/10.1016/j.uclim.2019.100454, 2019.
815	Thompson, G., Field, P. R., Rasmussen, R. M., and Hall, W. D.: Explicit Forecasts of Winter
816	Precipitation Using an Improved Bulk Microphysics Scheme. Part II: Implementation of a
817	New Snow Parameterization, Mon. Wea. Rev., 136, 5095–5115,
818	https://doi.org/10.1175/2008MWR2387.1, 2008.
819	Thundathil, R., Schwitalla, T., Behrendt, A., Muppa, S. K., ADAM, S., and Wulfmeyer, V.:
820	Assimilation of Lidar Water Vapour Mixing Ratio and Temperature Profiles into a
821	Convection-Permitting Model, JMSJ, https://doi.org/10.2151/jmsj.2020-049, 2020.
822	Thunis, P., Degraeuwe, B., Pisoni, E., Trombetti, M., Peduzzi, E., Belis, C. A., Wilson, J., and
823	Vignati, E.: Urban PM2.5 atlas: Air quality in European cities, JRC science for policy report,
824	28804, Publications Office, Luxembourg, 1 online resource, 2017.
825	UN: The World's Cities in 2018, United Nations, 2018.
826	Valcke, S., Balaji, V., Craig, A., DeLuca, C., Dunlap, R., Ford, R. W., Jacob, R., Larson, J.,
827	O'Kuinghttons, R., Riley, G. D., and Vertenstein, M.: Coupling technologies for Earth
828	System Modelling, Geosci. Model Dev., 5, 1589–1596, https://doi.org/10.5194/gmd-5-
829	1589-2012, 2012.
830	Vogel, B., Vogel, H., Bäumer, D., Bangert, M., Lundgren, K., Rinke, R., and Stanelle, T.: The
831	comprehensive model system COSMO-ART – Radiative impact of aerosol on the state of
832	the atmosphere on the regional scale, Atmos. Chem. Phys., 9, 8661–8680,
833	https://doi.org/10.5194/acp-9-8661-2009, 2009.





- 834 Wang, D., Stachlewska, I. S., Song, X., Heese, B., and Nemuc, A.: Variability of the Boundary
- 835 Layer Over an Urban Continental Site Based on 10 Years of Active Remote Sensing
- 836 Observations in Warsaw, Remote Sensing, 12, 340, https://doi.org/10.3390/rs12020340,
   837 2020.
- WHO: WHO Air quality guidelines for particulate matter, ozone, nitrogen dioxide and sulfur
  dioxide. Global update 2005., 2005.
- 840 Yang, Z.-L., Niu, G.-Y., Mitchell, K. E., Chen, F., Ek, M. B., Barlage, M., Longuevergne, L.,
- 841 Manning, K., Niyogi, D., Tewari, M., and Xia, Y.: The community Noah land surface model
- with multiparameterization options (Noah-MP): 2. Evaluation over global river basins, J.
  Geophys. Res., 116, https://doi.org/10.1029/2010JD015140, 2011.
- Zängl, G., Reinert, D., Rípodas, P., and Baldauf, M.: The ICON (ICOsahedral Non-hydrostatic)
  modelling framework of DWD and MPI-M: Description of the non-hydrostatic dynamical
  core, Q.J.R. Meteorol. Soc., 141, 563–579, https://doi.org/10.1002/qj.2378, 2015.
- Zhang, X., Huang, X.-Y., Liu, J., Poterjoy, J., Weng, Y., Zhang, F., and Wang, H.: Development
  of an Efficient Regional Four-Dimensional Variational Data Assimilation System for WRF,
  Journal of Atmospheric and Oceanic Technology, 31, 2777–2794,
- 850 https://doi.org/10.1175/JTECH-D-13-00076.1, 2014.
- Zhong, M., Saikawa, E., Liu, Y., Naik, V., Horowitz, L. W., Takigawa, M., Zhao, Y., Lin, N.-H.,
- and Stone, E. A.: Air quality modeling with WRF-Chem v3.5 in East Asia: sensitivity to
- emissions and evaluation of simulated air quality, Geosci. Model Dev., 9, 1201–1218,
  https://doi.org/10.5194/gmd-9-1201-2016, 2016.
- 855 Zieliński, M., Fortuniak, K., Pawlak, W., and Siedlecki, M.: Long-term Turbulent Sensible-
- Heat-Flux Measurements with a Large-Aperture Scintillometer in the Centre of Łódź,
  Central Poland, Boundary-Layer Meteorol, 167, 469–492,
- 858 https://doi.org/10.1007/s10546-017-0331-5, 2018.
- 859







Figure 1: Model domain 1 (a) and domain 3 (b). The blue dot in (a) denotes Stuttgart. Black dots in (b) show the location of the meteorological measurement sites. The blue diamond in (b) denotes the Neckartor (NT) location and the blue contour line denotes the Neckar River (River data © OpenStreetMap contributors 2020. Distributed under a Creative Commons BY-SA License).

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Figure 2: Land cover data from the Baden-Württemberg State Institute for the Environment, Survey and Nature Conservation (LUBW) reclassified for WRF in the innermost domain at a resolution of 50 m.

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Figure 3: NO₂ emissions valid at 07 UTC on January 21, 2019. (a) shows the emissions derived from the CAMS-REG-AP data set and (b) shows the emissions derived from the BW-EMISS data set (Map Data © OpenStreetMap contributors 2020. Distributed under a Creative Commons BY-SA License).

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Figure 4: Workflow of the AQFS prototype system.

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Bernhausen Ludwigsburg Stuttgart Am Neckartor Stuttgart Arnulf-Klett-Platz Stuttgart-Bad Cannstatt Figure 5: Observed NO₂ (a) and PM₁₀ (b) concentrations at several stations distributed over the model domain on 21 January 2019. The time zone (CET) corresponds to local time. Measurements at Neckartor, Hohenheimer Strasse, and Arnulf-Klett Platz are directly taken next to the main road.







Figure 6: (a) ECMWF operational analysis of 500 hPa geopotential height, sea level pressure (white contour lines) together with 500 hPa wind velocities valid at 00 UTC 21 January 2019. (b) shows the 925hPa equivalent potential temperature together with 925 hPa wind velocities and sea level pressure (white contour lines). Gray areas indicate values below the ECMWF model terrain. The black dot denotes Stuttgart and the reference wind vector length (top right corner of each Figure)) is equal to 25 m s⁻¹.

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Figure 7: Comparison of temperature, dewpoint and wind of the WRF model simulation (black line) and the sounding from Stuttgart-Schnarrenberg (red line) valid at 00 UTC (a) and 11 UTC (b) 21 January 2019. The solid lines denote the temperature profile and the dash-dotted line denotes the dewpoint profile. Wind barbs denote wind speed in m s⁻¹.

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Figure 8: 2-m temperature together with 10-m wind velocities at 12 UTC 21 January 2019. The thick black line denotes the Stuttgart city limits and the thin black contour lines denote the terrain. The blue line denotes the Neckar River (River data © OpenStreetMap contributors 2020. Distributed under a Creative Commons BY-SA License).

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Figure 10: Diurnal cycle of simulated sensible heat flux (SH, a) and ground heat flux (GRDFLX, b) at the three meteorological stations (white dots in Fig. 3). Positive values of GRDFLX indicate fluxes into the soil. The land cover categories are bare soil (airport), croplands (IPM), and urban (Schnarrenberg).

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Figure 11: Time-height cross section of the simulated potential temperature at IPM (top) and Schnarrenberg (bottom). The displayed altitude is above ground level (AGL).

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Figure 12: NO₂ concentration at the lowest model level for 07:30 UTC, 12 UTC, 18:00 UTC, and 23 UTC (from a to d) 21 January 2019. The black contour lines denote main roads and motorways in and around Stuttgart (Map Data © OpenStreetMap contributors 2020. Distributed under a Creative Commons BY-SA License). AP denotes the airport, A8 and A81 denote the main motorways around Stuttgart.

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Figure 13: Same as Fig. 12 but for PM₁₀ (Map Data © OpenStreetMap contributors 2020. Distributed under a Creative Commons BY-SA License).

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Figure 14: West-East cross section through Neckartor displaying the NO₂ concentration at 07:30 UTC (a) and 12 UTC (b), 21 January 2019. The red arrow denotes the motorway A81 and the white arrow denotes the Neckartor location. The black area shows the model terrain above mean sea level.

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 $Fig. \ 16: \ Time \ height \ cross \ section \ of \ NO_2 \ (top) \ and \ PM_{10} \ (bottom) \ at \ Neckartor \ (NT) \ up \ to \ an \ altitude \ of \ 450 \ m \ AGL.$