

Changes in biomass burning, wetland extent, or agriculture drive atmospheric NH₃ trends in key African regions

Jonathan E. Hickman^{1*}, Niels Andela^{2†}, Enrico Dammers³, Lieven Clarisse⁴, Pierre-François Coheur⁴, Martin Van Damme⁴, Courtney Di Vittorio⁵, Money Osohou⁶, Corinne Galy-Lacaux⁷, Kostas Tsigirdis^{1,8}, Susanne Bauer¹

¹NASA Goddard Institute for Space Studies, New York, USA

²NASA Goddard Space Flight Center, Beltsville, USA

³Air Quality Research Division, Environment and Climate Change Canada, Toronto, Canada

⁴Université libre de Bruxelles (ULB), Service de Chimie Quantique et Photophysique, Atmospheric Spectroscopy, Brussels, Belgium

⁵Wake Forest University, Winston-Salem, USA

⁶Laboratoire des Sciences de la Matière, de l'Environnement et de l'Energie Solaire, Université Félix Houphouët-Boigny, Abidjan, Côte d'Ivoire

⁷Laboratoire d'Aérodologie, Université Toulouse III Paul Sabatier / CNRS, France

⁸Columbia University, New York, USA

[†]Now at School of Earth and Ocean Sciences, Cardiff University, Cardiff, UK

*Correspondence to: jonathan.e.hickman@nasa.gov

Abstract

Atmospheric ammonia (NH₃) is a precursor to fine particulate matter and a source of nitrogen (N) deposition that can adversely affect ecosystem health. The main sources of NH₃—agriculture and biomass burning—are undergoing or expected to undergo substantial changes in Africa. Although evidence of increasing NH₃ over parts of Africa has been observed, the mechanisms behind these trends are not well understood. Here we use

31 observations of atmospheric NH₃ vertical column densities (VCDs) from the Infrared
32 Atmospheric Sounding Interferometer (IASI) along with other satellite observations of the
33 land surface and atmosphere to evaluate how NH₃ concentrations have changed over Africa
34 from 2008 through 2018, and what has caused those changes. In West Africa NH₃ VCDs are
35 observed to increase during the late dry season, with increases of over 6% yr⁻¹ in Nigeria
36 during February and March (p<0.01). These positive trends are associated with increasing
37 burned area and CO trends during these months, likely related to agricultural preparation.
38 Increases are also observed in the Lake Victoria Basin, where they are associated with
39 expanding agricultural area. In contrast, NH₃ VCDs declined over the Sudd wetlands in
40 South Sudan by over 1.5% yr⁻¹, though not significantly (p=0.28). Annual maxima in NH₃
41 VCDs in South Sudan occur during February through May and are associated with drying of
42 temporarily flooded wetland soils, which favor emissions of NH₃. The change in mean NH₃
43 VCDs over the Sudd is strongly correlated with variation in wetland extent in the Sudd: in
44 years when more area remained flooded during the dry season, NH₃ concentrations were
45 higher (r=0.64, p<0.05). Relationships between biomass burning and NH₃ may be observed
46 when evaluating national-scale statistics: countries with the highest rates of increasing NH₃
47 VCDs also had high rates of growth in CO VCDs; burned area displayed a similar pattern,
48 though not significantly. Livestock numbers were also higher in countries with
49 intermediate or high rates of NH₃ VCD growth. Fertilizer use in Africa is currently low but
50 growing; implementing practices that can limit NH₃ losses from fertilizer as agriculture is
51 intensified may help mitigate impacts on health and ecosystems.

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53 **1. Introduction:**

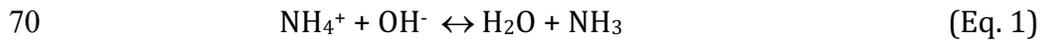
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55 Ammonia (NH₃), a reactive nitrogen (N) trace gas, plays a number of important roles
56 in the atmosphere, with implications for human health, climate, and ecosystems. Once in
57 the atmosphere, NH₃ contributes to the production of inorganic aerosols, the primary
58 constituents of fine particulate matter and a serious health hazard (Bauer et al., 2016;
59 Lelieveld et al., 2015; Pope et al., 2002). NH₃ can also be deposited to downwind
60 ecosystems, contributing to eutrophication, soil acidification, vegetation damage,
61 productivity declines, reductions in biodiversity, and indirect greenhouse gas emissions

62 (Denier Van Der Gon and Bleeker, 2005; Krupa, 2003; Matson et al., 1999; Stevens et al.,
63 2018; Tian and Niu, 2015).

64 Although NH₃ is emitted from natural soils, agriculture is by far the largest source of
65 NH₃ globally (Behera et al., 2013; Bouwman et al., 1997). Urea fertilizer and livestock
66 excreta are particularly important substrates for NH₃ formation, and can be volatilized
67 quickly under favorable environmental conditions (Bouwman et al., 1997). In all soils, NH₃
68 is formed in solution following the dissociation of ammonium (NH₄⁺; Eq. 1).

69



71

72 Soil NH₃ production is temperature-dependent, doubling with every 5°C temperature
73 increase, though the actual soil NH₃ flux is determined in part by plant and soil physiological
74 and physical factors (Sutton et al., 2013). On average, fertilizer use has been extremely low
75 in sub-Saharan Africa—often an order of magnitude or more lower than typical in Europe,
76 the United States, or China (Hazell and Wood, 2008; Vitousek et al., 2009). Livestock manure
77 N content also tends to be very low in sub-Saharan Africa (Rufino et al., 2006). The low
78 fertilizer use suggests that natural soils (as opposed to agricultural soils) may be a more
79 important source in the region than elsewhere in the world. However, agricultural
80 intensification and increasing fertilizer use has been a central policy focus for many African
81 countries, with national and regional efforts to increase N inputs by an order of magnitude
82 or more (AGRA, 2009).

83 After agriculture, biomass burning is the most important source of NH₃ globally
84 (Bouwman et al., 1997), with roughly 60 to 70% of global NH₃ emissions from fires occurring
85 in Africa (Cahoon et al., 1992; Whitburn et al., 2015). The amount of NH₃ emitted from
86 biomass fires is controlled primarily by the type of burning that occurs. N in fuel is present
87 predominantly in a chemically reduced state, and NH₃ is emitted in greater quantities from
88 low temperature smoldering combustion in which fuel N is incompletely oxidized (Goode et
89 al., 1999; Yokelson et al., 2008). Fuel moisture content, which can help determine whether
90 combustion is smoldering or flaming, is thus an important determinant of biomass burning
91 NH₃ emissions (Chen et al., 2010).

92 In contrast to other reactive N gases such as NO_x (nitric oxide + nitrogen dioxide),
93 NH₃ emissions are typically unregulated outside of Europe (Anker et al., 2018; Kanter,
94 2018; USDA Agricultural Air Quality Task Force, 2014), and substantial increasing trends
95 have been observed by the NASA Atmospheric InfraRed Sounder (AIRS) and the Infrared
96 Atmospheric Sounding Interferometer (IASI) over many of the world's major agricultural
97 and biomass burning regions during the 21st century (Van Damme et al., 2021; Warner et
98 al., 2017). West Africa has been identified as an important NH₃ source region (Van Damme
99 et al., 2018), where a trend of increasing NH₃ concentrations in recent decades has been
100 attributed at least in part to increased fertilizer use (Van Damme et al., 2021; Warner et al.,
101 2017). Increasing trends have also been observed over central Africa, and attributed to
102 higher rates of biomass burning (Van Damme et al., 2021; Warner et al., 2017). However,
103 the studies by Warner et al. (2017) and Van Damme et al. (2021) were global in nature, and
104 as such could not include detailed explorations of the drivers of trends such as
105 consideration of emission seasonality or the geographic distribution of emission drivers.
106 Consideration of these factors is particularly important across large parts of Africa where
107 both biomass burning and soils are potentially important sources of NH₃ (van der A et al.,
108 2008).

109 Here we use an eleven-year satellite record to evaluate trends in atmospheric NH₃
110 concentrations over Africa from 2008 through 2018, including detailed examination of
111 three regions where changes are pronounced: West Africa, the Lake Victoria Region, and
112 South Sudan.

113

114 **2. Data and Methods**

115 **2.1 Global gridded data**

116 Multiple data products were used, including satellite observations and spatial
117 datasets:

118 -IASI-A, launched aboard the European Space Agency's MetOp-A in 2006, provides
119 measurements of atmospheric NH₃ and carbon monoxide (CO) twice a day (9:30 in the
120 morning and evening, Local Solar Time at the equator). Here we use morning observations,
121 when the thermal contrast is more favorable for retrievals (Clarisse et al., 2009; Van
122 Damme et al., 2014a). The NH₃ retrieval product used (ANNI-NH₃-v3R) follows a neural

123 network retrieval approach. We refer to Van Damme et al. (2017) and Van Damme et al.
124 (2021) for a detailed description of the algorithm. For CO, we used the product obtained
125 with the FORLI v20140922 retrieval algorithm (Hurtmans et al., 2012). Given the absence
126 of hourly or even daily observations of NH₃ concentrations in sub-Saharan Africa, the
127 detection limit of IASI is difficult to determine with certainty. However, the region
128 experiences high thermal contrast, and IASI seems to be able to reliably observe NH₃ down
129 to 1 to 2 ppb at the surface (Clarisse et al., 2009; Van Damme et al., 2014b). We gridded the
130 Level-2 IASI NH₃ and CO products to 0.5° × 0.5° resolution. We used a conventional binning
131 approach based on the center of each satellite footprint. We did not apply an averaging
132 weight. Quality control procedures were followed as detailed in van Damme et al. 2017 and
133 Van Damme et al., 2021. Specifically, the screening of retrievals included filtering of
134 retrievals where cloud cover is over 10%, where the total column density is below zero and
135 the absolute value of the hyperspectral range index (HRI) is above 1.5, and where the ratio
136 of the total column density to HRI is larger than 1.5×10^{16} molecules cm⁻².

137 The IASI products have been validated using ground-based Fourier transform
138 infrared (FTIR) observations of NH₃ total columns, with robust correlations at sites with
139 high NH₃ concentrations, but lower at sites where atmospheric concentrations approach
140 IASI's detection limits (Dammers et al., 2016; Guo et al., 2021). Compared to the FTIR
141 observations, total columns from previous IASI NH₃ products (IASI-LUT and IASI-NNv1)
142 are biased low by ~30% which varies per region depending on the local concentrations.
143 Although FTIR observations are absent from Africa, earlier work has shown fair agreement
144 between previous versions of IASI total column densities and surface observations of NH₃
145 using passive samplers across the International Network to study Deposition and
146 Atmospheric chemistry in AFrica (INDAAF) network in West Africa (Van Damme et al.,
147 2015), including in observations of seasonal variation (Hickman et al., 2018; Ossouhou et al.,
148 2019). Validation of the IASI CO product using surface, aircraft, and satellite observations
149 have found total columns to have an error that is generally below 10-15% in the tropics
150 and mid-latitudes (George et al., 2009; Kerzenmacher et al., 2012; Pommier et al., 2010; De
151 Wachter et al., 2012). The IASI NH₃ and CO products were used for the years 2008—the
152 first full year of data available—to the end of 2018. Random errors in observations can be

153 assumed to cancel out in the annual mean, which is what we used in our analysis. With the
154 assumption that random errors cancel out, only systematic errors related to tropospheric
155 vertical column contents remain; these systematic errors do not contribute to uncertainty
156 in trend analyses. In addition, we first take monthly averages based on all daily
157 observations within a given month before calculating seasonal means to minimize any
158 potential effects of temporal variability in cloud cover.

159 -The Tropical Rainfall Measuring Mission (TRMM) daily precipitation product (3B42)
160 is based on a combination of TRMM observations, geo-synchronous infrared observations,
161 and rain gauge observations (Huffman et al., 2007). Independent rain gauge observations
162 from West Africa have been used to validate the product, with no indication of bias in the
163 product (Nicholson et al., 2003).

164 - NOAA Global Surface Temperature Dataset, a 0.5° gridded 2m monthly land surface
165 temperature product (Fan and van den Dool, 2008). The data set is based on a combination
166 of station observations from the Global Historical Climatology Network version 2 and the
167 Climate Anomaly Monitoring System (GHCN_CAMS), and uses an anomaly interpolation
168 approach which relies on observation-based reanalysis data to derive spatio-temporal
169 variation in temperature lapse rates for topographic temperature adjustment.

170 - 500m MCD64A1 collection 6 Moderate Resolution Imaging Spectroradiometer
171 (MODIS) burned area product for the period 2008-2018 (Giglio et al., 2018). The burned
172 area data are aggregated by month and gridded to 0.25° resolution, and do not include
173 burned area from small fires.

174 - MODIS MCD12C1 (collection 5) land cover product, which provides the percentage
175 of cropped area in each 0.25° grid cell (Friedl et al., 2002). In Africa, agriculture is often
176 practiced in complex mosaics of agricultural and natural land cover, so we used both the crop
177 and crop/natural area mosaic MODIS classifications as agricultural area in our analysis.

178 -We also used data on the spatio-temporal distribution of armed conflict events from
179 the Armed Conflict Location & Event Data Project (ACLED; Raleigh et al., 2010). We included
180 data for both violent and non-violent conflict events over the period 2008-2018.

181 **2.2 Sudd wetland extent**

182 Monthly flooded area extents of the Sudd Wetland, South Sudan from 2000 to 2017
183 were derived from 8-day composite MODIS land surface reflectance imagery (MOD09A1);
184 data from 2005 through 2017 were used in the analyses. We refer to Di Vittorio and
185 Georgakakos (2018) for a detailed description of the classification procedure designed to
186 retrieve these data. In summary, monthly flood maps were obtained through a two-stage
187 classification procedure. The first stage used the full 18-year data set to produce a wetland
188 land cover map that distinguishes between wetland vegetation classes and their flooding
189 regimes (permanently flooded, seasonally flooded, or non-flooded). The second stage
190 compares seasonally flooded pixels from each vegetation class to their non-flooded
191 counterparts on a monthly basis to identify the timing and duration of flooding for each pixel.
192 These data were originally derived to calibrate a hydrologic model of the Sudd that is
193 dependent on Nile flows; therefore, a connectivity algorithm was applied to ensure that all
194 flooded pixels were physically connected to the Nile River. A few adjustments have been
195 made to the previously published dataset for the application of this study. The classification
196 algorithm has been improved to more accurately capture the inter-annual fluctuations in the
197 permanently flooded areas. The dataset was also extended through 2017, and the total
198 flooded area was quantified prior to applying the connectivity algorithm. The magnitudes of
199 the monthly flooded area estimates are now substantially larger because they include areas
200 flooded from local runoff in addition to areas flooded by the Nile River.

201

202 **2.3 Spatial and national analyses**

203 We evaluated spatial relationships between mean annual tropospheric NH₃
204 concentration and several independent variables at 0.25° resolution: population density,
205 livestock density, and cropped area. Population density and livestock density data are not
206 available as time series suitable for trend analysis, so we use single year values in our
207 analyses. We calculated population density based on the 2017 version of the US
208 Department of Energy's Gridded Landscan population dataset (Dobson et al., 2000;

209 available at <https://landscan.ornl.gov>). Livestock density was based on the FAO global
210 gridded livestock dataset for the year 2007 (Robinson et al., 2014). Before analysis, we
211 converted the livestock densities of chickens, goats, pigs, and sheep to tropical livestock
212 units (TLU), using values of 0.01, 0.1, 0.2, and 0.1 TLU, respectively; North African cattle
213 were converted using a factor of 0.7, whereas sub-Saharan cattle were converted using a
214 factor of 0.5 (Chilonda and Otte, 2006). For cropped area, we used the MODIS MCD12C1
215 (collection 5) land cover product as described above. We conducted spatial analyses by
216 establishing a map of 3° grid cells and calculating the correlation between the value of each
217 independent variable and NH₃ for all 0.5° grid cells within the larger grid cells (N = 36
218 including water grid cells, though these were excluded from the analysis).

219

220 National data on annual livestock numbers, crop production, and fertilizer N use were
221 obtained from the UN Food and Agriculture Organization FAOSTAT for 51 African countries
222 (FAO, 2020). Livestock data consisting of sheep, goats, cattle, and pigs were converted to
223 tropical livestock units as described above, and buffaloes were converted using a conversion
224 factors of 0.7 (Chilonda and Otte, 2006). National emissions of CO₂ were obtained from
225 World Bank Open Data (World Bank, 2019). National-level mean annual cropland area,
226 burned area, and atmospheric NH₃ and CO concentrations were also calculated for each of
227 the 51 countries from the spatial datasets described above. Countries were sorted into three
228 bins based on whether their relative change in mean annual NH₃ concentration was low,
229 medium or high, and means and standard errors were calculated for each of the three 17-
230 country bins.

231 Linear trend analyses were conducted using linregress from the scipy.stats package
232 in Python v3.6.3. Statistical analyses of national scale data were conducted using ANOVA in
233 R. Data were log or rank transformed when necessary to meet the assumptions of ANOVA.
234 Values of α for treatment comparisons following significant ANOVA results were corrected
235 for multiple testing using Benjamini-Hochberg corrections.

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238 **3. Results & Discussion**

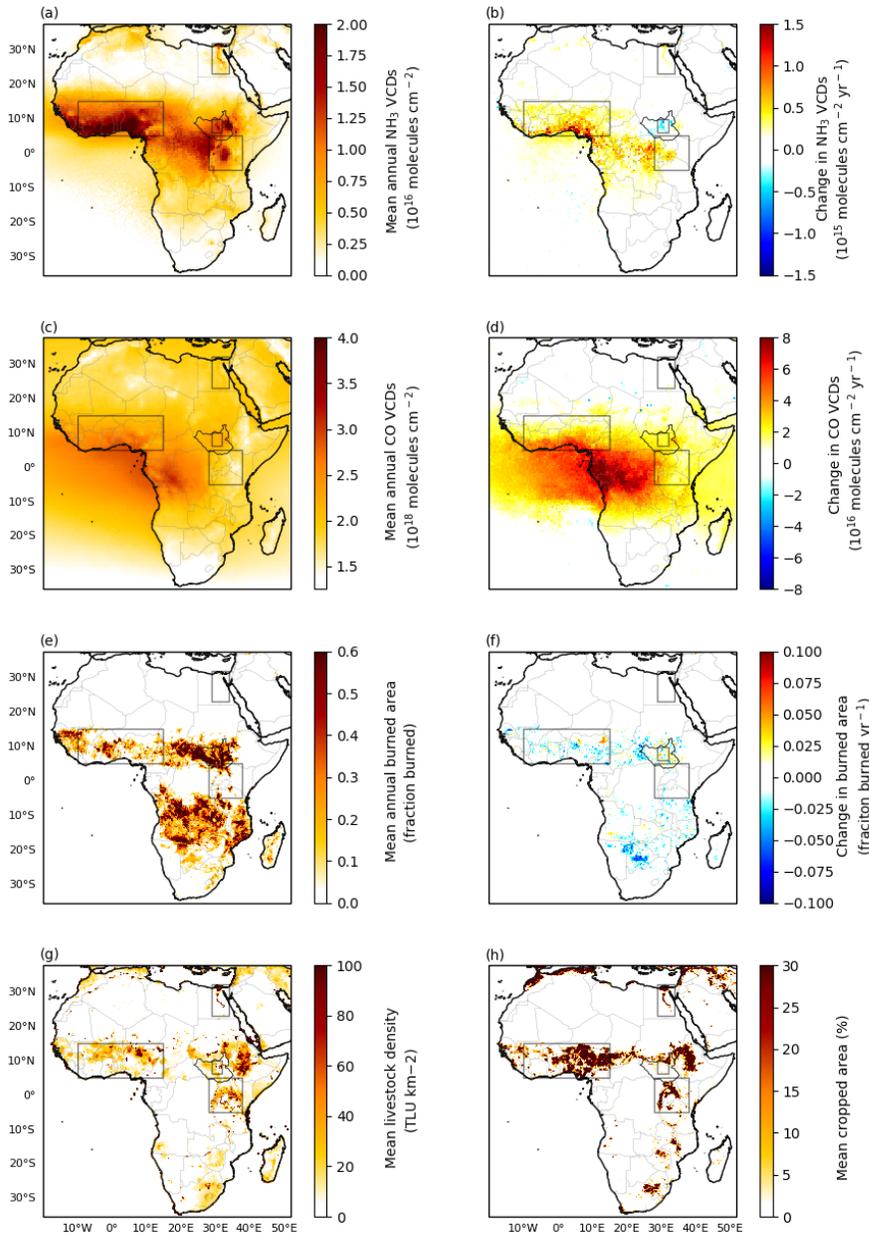
239 3.1 Continental distributions and trends

240 Mean annual NH₃ concentrations for 2008-2018 are highest across the savannas
241 and forest-savanna mosaics in North equatorial Africa, and especially in West Africa; there
242 are smaller regional hotspots in the Lake Victoria basin, South Sudanese wetlands, and
243 along the Nile delta and river (Fig. 1a). Parts of these regions experience substantial
244 biomass burning (Fig. 1e), high livestock densities (Fig. 1g), and/or high cropland cover
245 (Fig. 1h), all of which can contribute to NH₃ emissions. The high concentrations in West
246 Africa, which is one of the major global NH₃ hotspots (Van Damme et al., 2018), is likely the
247 result of biomass burning emissions. Biomass burning emissions tend to drive seasonal
248 variation in NH₃ VCDs in West Africa, with the largest emissions occurring late in the dry
249 season and early rainy season (Hickman et al., 2021b). In addition to local emissions,
250 biomass burning emissions and their reactive products are transported to the coast of West
251 Africa during both the northern hemisphere rainy season, when it is transported from
252 central and southern Africa, and during the dry season, when it is transported from
253 biomass burning regions to the east (Sauvage et al., 2007). Most areas with trends are
254 significant at P=0.2 or higher (Fig. S1).

255 In addition to being hotspots of mean NH₃ concentrations, some of these regions
256 have also experienced increases in NH₃ concentrations from 2008 to 2018 (Fig. 1b). Like
257 Warner et al. (2017) and Van Damme et al. (2021), we observed some increases in the
258 northern grasslands, central African forests, and the Nile region, but we also observe trends
259 in the Lake Victoria Basin, which Warner et al. (2017) did not, but Van Damme et al. (2021)
260 did. Also in contrast to Warner et al. (2017) but in line with Van Damme et al. (2021), we
261 observe a prominent decline in NH₃ VCDs over South Sudan (Fig. 1b, S1).

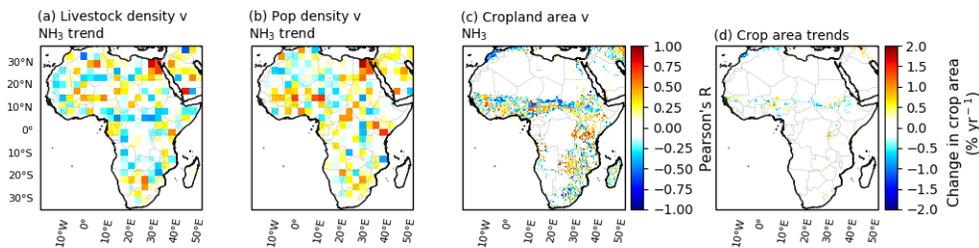
262 The Nile region exhibits elevated NH₃ concentrations and a modest positive trend
263 over the observation period (Fig. 1a, 1b). This trend appears largely to be related to
264 agriculture and livestock: in a spatial analysis, snapshots of livestock densities and of
265 population densities are both positively related to changes in NH₃ VCDs (Fig. 2). Although
266 there is not a positive relationship between agricultural area and NH₃ VCDs over the Nile
267 region from 2008 to 2018, Egypt's population increased by roughly 25% over that period
268 (World Bank, 2019), and fertilizer N use increased by roughly 8% after a decline in use

269 between 2004 and 2007 (FAO, 2020), suggesting that increased agricultural N inputs may
270 be contributing to the trend. We evaluate the other regions in more detail below.



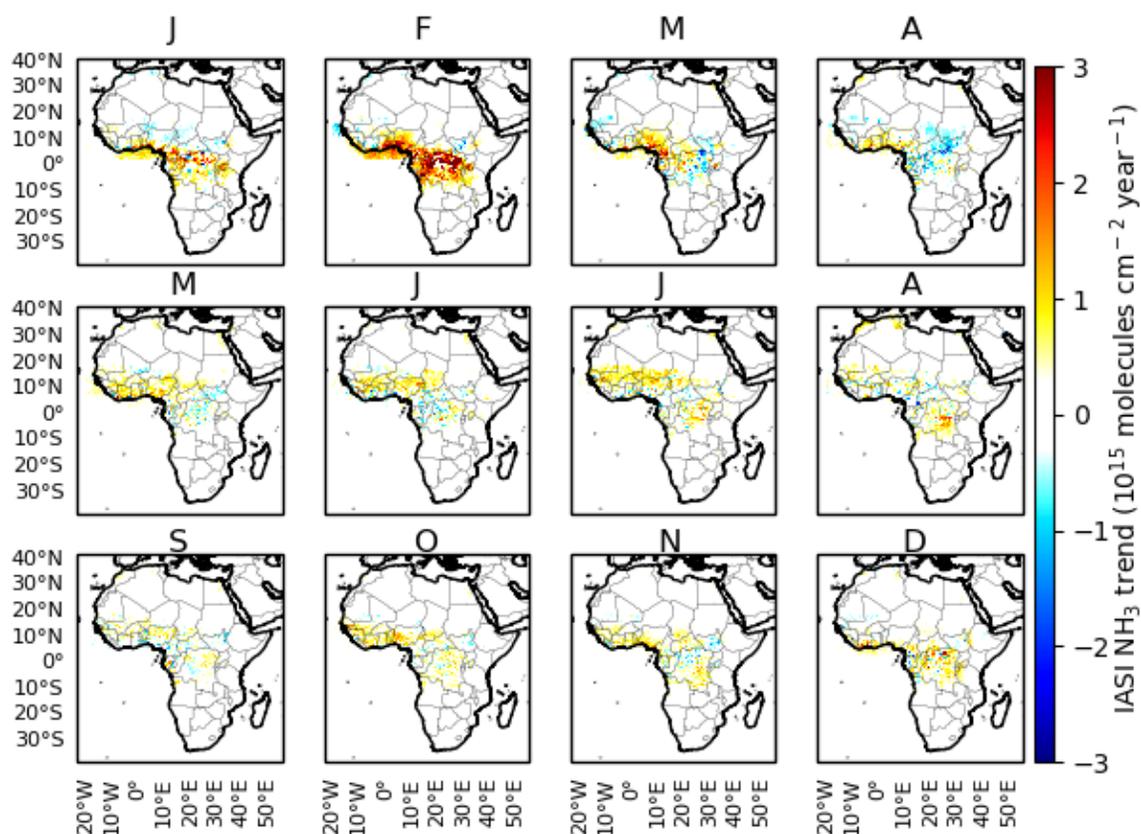
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272 **Figure 1.** Annual averages and trends in atmospheric NH₃ VCDs, CO VCDs, and burned
273 area, as well as spatial distribution of livestock density and cropped area across seven sub-
274 Saharan African ecoregions. Mean annual (a) and trend (b) in atmospheric NH₃ VCDs from

275 IASI for the period 2008 through 2018. Mean annual (c) and trend (d) in annual
 276 atmospheric CO VCDs from IASI for the same period. Mean annual (e) and trend (f) in
 277 annual burned area from MODIS for 2008-2018. Livestock densities for 2007 from the FAO
 278 (f), and mean cropped area from MODIS for 2008-2018 (g). The border of South Sudan is
 279 highlighted in black, and several regions boxed: the Nile region at 30°N, the Sudd wetland
 280 in South Sudan, the Lake Victoria region at the equator, and West Africa centered around
 281 10°N.



283 **Figure 2.** Relationships between NH₃ trends and livestock density, population density, and
 284 cropland area, as well as changes in cropland area. Spatial correlations between changes in
 285 annual atmospheric NH₃ VCDs and livestock density (a) and population density (b).
 286 Correlation between cropland area and NH₃ VCDs for 2008 through 2018 (c). Change in
 287 cropland area for 2008 through 2018 (d). The NH₃ and crop area trends are based on data for
 288 2008 through 2018, livestock density data are for the year 2007, population density data
 289 are for the year 2017.

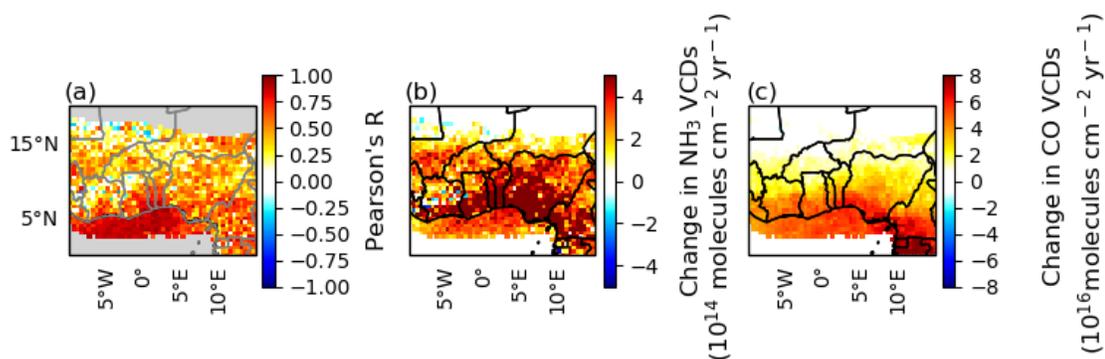
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292 **Figure 3.** Change in mean monthly atmospheric NH₃ VCDs for the period 2008 through
 293 2018. Grid cells where mean annual NH₃ VCDs for the entire period are under 5x10¹⁵
 294 molecules cm⁻² are not displayed. Results significant at P=0.05 are presented in Figure S5.

295



296

297

298 **Figure 4.** Correlation coefficient for the relationship between mean annual CO and
299 NH₃ VCDs (a), changes in NH₃ VCDs (b) and changes in CO VCDs (c) over 2008 through
300 2018 in West Africa. Grid cells where mean annual NH₃ VCDs for the entire period are
301 under 5×10^{15} molecules cm⁻² are not displayed. Results significant at P=0.05 for the entire
302 continent are presented in Figure S6.

303

304 **3.2 West Africa**

305 The increasing trend in NH₃ VCDs over West Africa are centered over Nigeria and
306 the southern coast, and to a lesser extent across parts of the wet savanna (Fig. 1b).
307 Increases in NH₃ VCDs tend to be higher in grid cells with higher population densities in
308 Nigeria and other parts of West Africa (Fig. 2b), suggesting a possible anthropogenic
309 influence. The spatial distribution of the mean annual NH₃ trend is overlapped by a
310 substantial increase in mean annual CO VCDs (Fig. 1b, 1d), pointing to a biomass burning
311 source, as is also the case in central Africa. Earlier studies have found substantial declines
312 in annual burned area across the north equatorial African biomass burning region as
313 detected by MODIS (Andela et al., 2017; Andela and van der Werf, 2014) and related
314 declines in NO₂ VCDs across the region (Hickman et al., 2021a), which would seem to stand
315 in contrast to the increasing CO and NH₃ trends observed here.

316 However, the annual decline in burned area and NO₂ VCDs is characterized by
317 heterogeneity when considering individual months. In West Africa, the dry season is
318 typically November to February or March. During the transition from the dry to rainy
319 season in February and March, NO₂ VCDs exhibit increasing rather than decreasing trends
320 in West Africa, though burned area patterns are not as clear when 2018 is included
321 (Hickman et al., 2021a; Fig. S2, S3). Although these increases in NO₂ VCDs are small in the
322 annual context, they occur at a time of year when biomass burning combustion is less
323 complete, potentially due to greater fuel moisture and declining fire radiative power
324 (Hickman et al., 2021a; Zheng et al., 2018). These conditions would lead to greater

325 emissions of less oxidized species such as CO and NH₃, rather than the more fully oxidized
326 species such as CO₂ and NO₂ that dominate emissions during the peak of the biomass
327 burning season (Fig. S2, S4). Indeed, our observations suggest that much of the increasing
328 NH₃ trend occurs during this transitional period, with NH₃ VCDs increasing by roughly 6%
329 yr⁻¹ for all of Nigeria during February and March (Fig. 3, S5; p<0.01). Variation in NH₃ VCDs
330 are positively correlated with CO VCDs (Fig. 4a, S6), which are also increasing during this
331 period (Fig. 4c, S4).

332 These correlations imply a biomass burning source for the increasing NH₃ VCDs in
333 West Africa; although the burned area trends are not as clear, it is important to remember
334 that MODIS undercounts burned area during this time of year by a factor of 3 to 6, and so
335 would be less sensitive to trends (Ramo et al., 2021; Roteta et al., 2019). Although there is
336 considerable gas flaring in Nigeria, gas flaring emissions have exhibited long-term negative
337 trends (Doumbia et al., 2019). In addition, although NO₂ VCDs were found to decrease
338 across the productive savannas of West Africa, regions of increasing NO₂ VCDs were
339 observed over large parts of Nigeria, further suggesting that there may be increases—or at
340 least smaller decreases—in biomass burning in the country (Hickman et al., 2021a). It is
341 unlikely that changes in chemical sinks—specifically, the formation of nitrate aerosols in
342 reactions with NO_x or sulfate—are responsible for the increasing trend: the observed
343 increase in NO₂ VCDs observed during February and March would be expected to lead to a
344 shorter NH₃ lifetime and decreasing VCDs. In addition, emissions of SO₂ are relatively low
345 in West Africa, with moderate emissions occurring in Nigeria, but neither emissions nor
346 lifetime exhibit clear seasonal variation (Lee et al., 2011).

347 Small agricultural fires are likely an important contributor to the increasing NH₃
348 VCDs during the dry-to-rainy season transitional period—a period when agricultural fires
349 are common in the region (Korontzi et al., 2006). There are large numbers of small fires
350 that are not detected by MODIS during these months: as noted above, estimates of burned
351 area during February, March, and April are revised upwards by roughly a factor of 3 to 6
352 over MODIS when small fires are included (Ramo et al., 2021; Roteta et al., 2019). Many of
353 these small fires are likely related to agricultural field preparation prior to planting
354 (Gbadegesin and Olusesi, 1994), which typically takes place in March or April (Vrieling et
355 al., 2011; Yegbemey et al., 2014). An increase in fires during this transitional period is also

356 consistent with one of the primary mechanisms behind the overall decline in burned area:
357 roughly half of the decline is attributed to increased population density and the expansion
358 of agricultural area, which contributes to the anthropogenic suppression of larger fires
359 (Andela et al., 2017; Andela and van der Werf, 2014). This agricultural expansion,
360 however, can be expected to be accompanied by increases in small fires used for the
361 removal of stubble or harvest byproduct (Gbadegesin and Olusesi, 1994), leading to the
362 increased emissions during the rainy-to-dry season transition observed here.

363 Globally, agricultural emissions from fertilized soils and livestock excreta are the
364 largest source of NH_3 (Bauer et al., 2016), and Warner et al. (Warner et al., 2017) suggest
365 that national-scale changes in fertilizer use could explain the NH_3 trend over Nigeria.
366 However, as noted above, much of the increase in West Africa occurs prior to the start of
367 the planting season—before fertilizer is applied—and appears likely to be due to biomass
368 burning emissions instead, potentially related to field preparation. Fertilizer or manure
369 may make a contribution to the increasing trend later in the year, as NH_3 VCDs increase in
370 the wet savanna during May, June, and July (Fig. 3), though there are also significant
371 correlations between NH_3 and CO VCDs (Fig. 4), suggesting that biomass burning may
372 continue to play an important role. However, average N fertilizer use in West Africa is
373 universally under $40 \text{ kg N ha}^{-1} \text{ yr}^{-1}$, typically under $20 \text{ kg N ha}^{-1} \text{ yr}^{-1}$, and is under 10 kg N
374 $\text{ha}^{-1} \text{ yr}^{-1}$ in Nigeria—over an order of magnitude lower than rates in Europe, the United
375 States, and China (FAO, 2020). Although percentage changes in fertilizer use are
376 substantial, in absolute terms they represent increases of less than $2 \text{ kg N ha}^{-1} \text{ yr}^{-1}$, and
377 frequently less than $1 \text{ kg N ha}^{-1} \text{ yr}^{-1}$, a relatively small but perhaps not entirely trivial
378 perturbation to the N cycle: Between 2000 and 2007, total N deposition averaged 8.38 kg N
379 $\text{ha}^{-1} \text{ yr}^{-1}$ in wet savanna and $14.75 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ in forest ecosystems based on surface
380 sampling sites (Galy-Lacaux and Delon, 2014), and biological N fixation in tropical and wet
381 savannas has been estimated as ranging from 16 to $44 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ (Bustamante et al.,
382 2006). These estimates suggest that fertilizer increases may represent a 1 to 2% annual
383 increase in N inputs. But given the small magnitude of fertilizer applications, it appears
384 unlikely that changes in fertilizer use can explain the entirety of NH_3 increases during the
385 growing season. Our analyses do suggest that livestock may contribute to increasing NH_3
386 VCDs over the Sahel, from roughly 15 to 18N (Fig. 2a). However, many of these pixels are

387 also those where population density appears to be playing a role (Figure 2b) and where
388 correlations between NH_3 and CO VCDs are present during the transition from the dry to
389 rainy season (Fig. S7), which may reflect a contribution from agricultural fires.

390

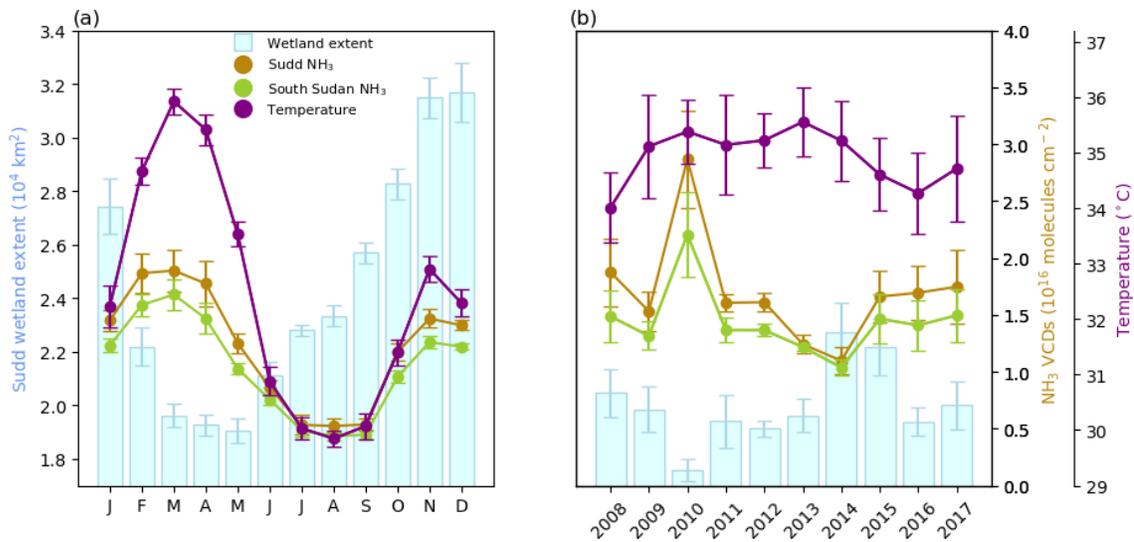
391 **3.3. South Sudan**

392 The most notable declining trend in NH_3 VCDs occurs in South Sudan over the Sudd
393 wetlands at a rate of over $2\% \text{ yr}^{-1}$ (Fig. 1b; $p=0.20$). It appears that this decline is related to
394 interannual variation in the flooded extent of the Sudd, a vast wetland that connects the
395 White and Blue Nile tributaries. Seasonal variation of inflow to the Sudd leads to variation
396 in flooded extent: an area of roughly $15,000 \text{ km}^2$ is permanently flooded, and another
397 roughly $15,000 \text{ km}^2$ is temporarily flooded each year, with considerable interannual
398 variation in the total flooded area (Di Vittorio and Georgakakos, 2018). Among other
399 factors, drying soils should increase production and emissions of NH_3 from soils, as Eq. 1 is
400 shifted to the right (Clarisse et al., 2019). Earlier work evaluating an NH_3 hotspot over Lake
401 Natron in Tanzania found that the drying of seasonally flooded soils leads to large
402 emissions of NH_3 : As the waters of Lake Natron recede during the dry season each year
403 and the surrounding mud flats dry out, NH_3 VCDs increase rapidly, with hotspots appearing
404 over the mudflats (Clarisse et al., 2019). These elevated VCDs are attributed to multiple
405 possible factors, including the effects of drying on concentrations of NH_3 in solution (which
406 increases the concentration gradient with the atmosphere), reduced biological uptake of
407 NH_3 , convective transport of dissolved NH_3 from depth to the soil surface, and increased
408 mineralization of labile organic matter (Clarisse et al., 2019).

409 We find the same clear seasonal relationship between wetland flooded extent and
410 NH_3 concentrations over the Sudd—VCDs increase as waters recede from the temporarily
411 flooded area, leading to annual maxima from February through May (Fig. 5a; bounding box
412 of 29E to 31.5E and 6N to 9.9N). Like the entire country, seasonal variation in NH_3 VCDs
413 over the Sudd follow variation in surface temperature, but NH_3 concentrations over the
414 Sudd are substantially elevated compared to surrounding regions during this time of year
415 but not others, suggesting that a mechanism in addition to temperature is contributing to

416 the elevated emissions in the Sudd during February through May, a period that spans the
 417 end of the dry season and start of the rainy season (Fig. S8). This conclusion is supported
 418 by an analysis of interannual variation of VCDs during the February through May period:
 419 Interannual variation in NH₃ VCDs is largely decoupled from variation in temperature, but
 420 NH₃ VCDs appear to vary inversely with the amount of area that dries out each year (Fig.
 421 5b). Over the period for which flooded extent data are currently available for the Sudd, the
 422 minimum flooded extent tends to increase—that is, less area dries out each year—resulting
 423 in an overall decline in NH₃ VCDs. Linear regression reveals that this change in flooded
 424 extent explains a large proportion of the annual variation in NH₃ in the Sudd bounding box
 425 ($r=-0.64$, $p=0.046$), as well as for the country as a whole ($r=-0.60$, $p=0.065$). These analyses
 426 strongly suggest that the declining trend in NH₃ over the Sudd is a direct result of an overall
 427 increase in the minimum flooded extent over the observation period.

428
 429



430

431 **Figure 5.** Mean (a) monthly and (b) February through May annual mean flooded
 432 extent of the Sudd, surface temperatures over South Sudan, and NH₃ VCDs over the Sudd
 433 and the entirety of South Sudan for the period 2008 through 2017.

434

435
436

437 It is possible that conflict in South Sudan could contribute to the decline in NH₃
438 VCDs. In 2013, a civil conflict emerged in South Sudan that was ultimately responsible for
439 the displacement of millions of people (Global Internal Displacement Monitoring Centre,
440 2020; World Bank, 2019) and the disruption of livestock migration patterns (Idris, 2018).
441 However, these disruptions appeared only after the onset of the long-term change in NH₃,
442 and appear unlikely to make an important contribution to the observed interannual
443 variation (SI Text, Fig. S9, S10).

444 It is unlikely that changes in chemical sinks are responsible for the decline in NH₃
445 VCDs. VCDs of tropospheric NO₂ are also decreasing in the region (Fig. S11), which is
446 suggestive of less formation of particulate-phase ammonium rather than more.
447 Anthropogenic SO₂ emissions in Africa in general and South Sudan in particular are very
448 low (European Commission Joint Research Centre (JRC)/Netherlands Environmental
449 Assessment Agency (PBL), 2016), and would not be expected to be emitted from the Sudd;
450 more generally, the clear spatial association between the NH₃ trend and the Sudd (Fig. 1,
451 Fig. S12) is strongly suggestive of changes in emissions rather than atmospheric processes
452 being responsible for the trend.

453

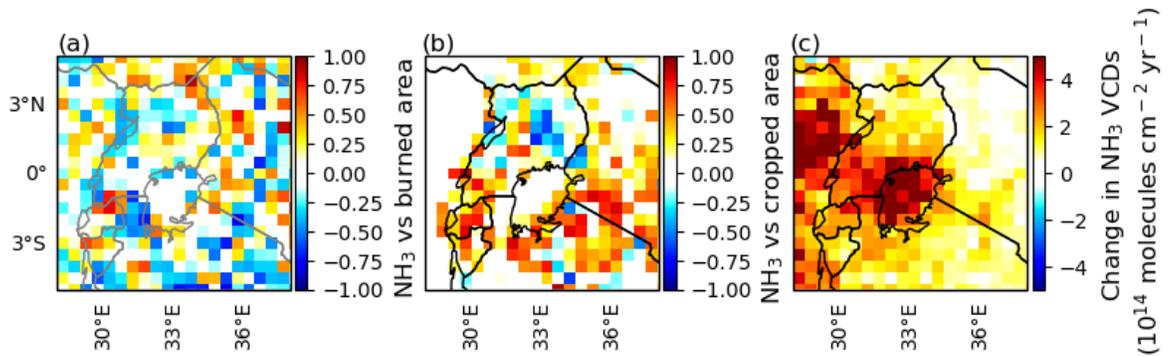
454 **3.4 Lake Victoria Basin region**

455 The Lake Victoria Basin and its surroundings—an area including elevated mean NH₃
456 VCDs—exhibit an increasing NH₃ trend (Fig. 1b, Fig. 7, Fig. S13), which appears to be the
457 result of increasing agricultural activity in the area. The region includes a high and
458 increasing density of agricultural land (Fig. 1h, Fig. 2d, Fig. S14), and these increases in
459 cropped area are positively correlated with increases in NH₃ VCDs across much of the
460 region (Fig. 2c). The northern and southern halves of the Lake Victoria region—which
461 straddles the equator—have distinct growing seasons: in the north, the season generally

462 starts in April, whereas in the south, it starts in November or December (Vrieling et al.,
463 2011). Some of the long-term trend reflects this seasonality, with increases in the north
464 and south occurring during their respective growing seasons (Fig. 3, Fig. S15). Fertilizer
465 use in the Lake Victoria region is low: national averages range from about 1 to 3 kg
466 nutrients ha⁻¹ yr⁻¹ in Uganda to about 35 to 40 kg nutrients ha⁻¹ yr⁻¹ in Kenya (Elrys et al.,
467 2019; World Bank, 2019); to put these numbers in context, Organization for Economic
468 Cooperation and Development (OECD) countries use about 135-140 kg nutrients ha⁻¹ yr⁻¹
469 (World Bank, 2019). Although rates of fertilizer use have increased by substantial
470 proportions, the absolute amount of increase is relatively small, typically roughly 1 to 10 kg
471 nutrients decade⁻¹. Unlike in West Africa, however, interannual variation in burned area
472 (Fig. 7, S16) does not exhibit a clear relationship with changes in NH₃ VCDs. Consequently,
473 we expect that both the expansion and intensification of agriculture in the region
474 contribute to the increasing NH₃ VCDs.

475 We note that there is a negative correlation between cropland area and NH₃ VCDs in
476 Uganda, north of Lake Victoria (Figure 7b). We expect this is a consequence of the
477 extremely low fertilizer use in Uganda (Masso et al., 2017), which leads to depletion of soil
478 N—and thus substrate for ammonia volatilization—over time (Cobo et al., 2010).

479 We also note that there is an apparent increase in NH₃ VCDs over the lake itself. It is
480 important to note that differences in conditions over the lake and adjacent land cover—e.g.,
481 emissivity, thermal contrast, etc.—contribute to substantial differences in mean retrieved
482 NH₃ VCDs over the lake relative to the surrounding land surface. Both monthly and
483 interannual variation in NH₃ VCDs over Lake Victoria correspond closely to variation in
484 NH₃ VCDs over the surrounding land surface (Figure S17, S18), suggesting that the trend
485 over the lake results from transport of NH₃ emitted from the surrounding land surface.



486

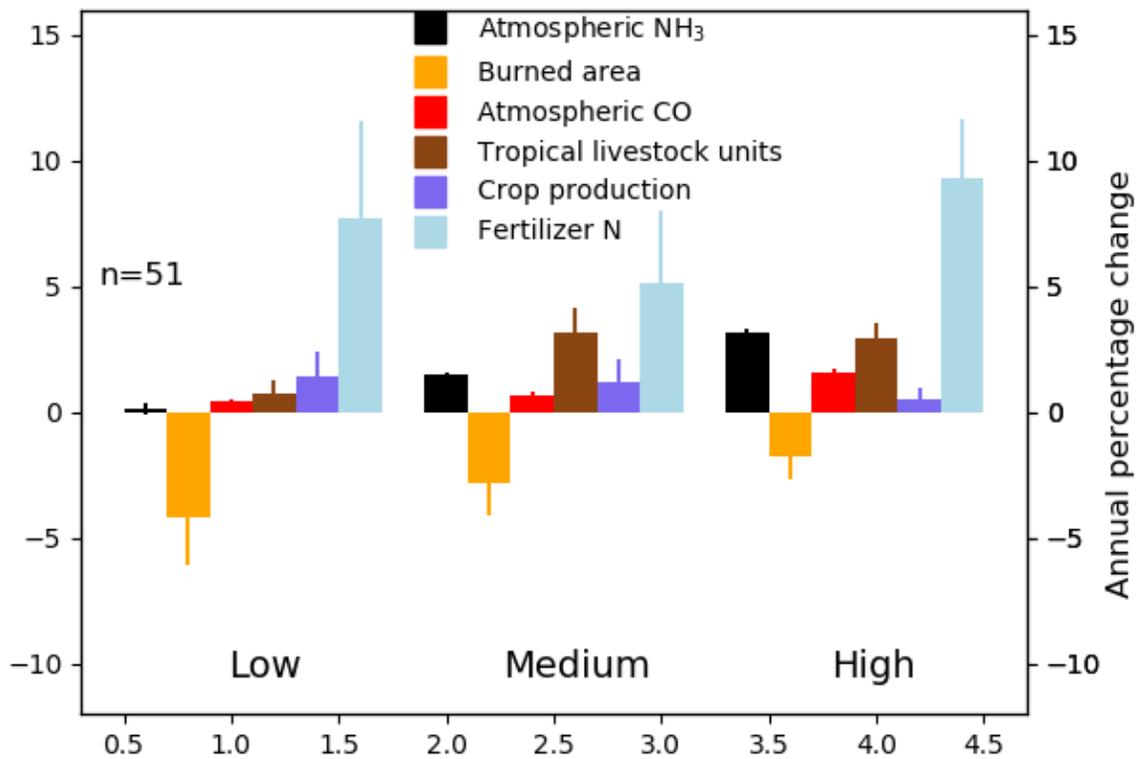
487 **Figure 7. Changes in NH₃ VCDs and their relationship with burned area and**
 488 **cropped area** over the Lake Victoria region for the 2008 through 2018 period. (a)
 489 Correlation coefficients for the relationship between NH₃ VCDs and burned area. (b)
 490 Correlation coefficients for the relationship between NH₃ VCDs and cropped area, including
 491 mosaics of crops and natural vegetation cover. (c) Changes in NH₃ VCDs.

492

493 3.5 National-scale relationships

494 Examining relationships at a national scale can provide insight into relationships
 495 between changes in agricultural or biomass burning and changes in atmospheric NH₃ VCDs
 496 at larger scales. When grouping countries into three bins based on their annual percentage
 497 changes in NH₃ VCDs, there is some evidence for a broad relationship between livestock
 498 and NH₃ VCDs at the national scale (Fig. 8). The rate of change in national-scale NH₃ VCDs
 499 varies significantly among bins ($p < 0.001$; rank transformed, though note that residuals
 500 may still deviate from normality). The annual percentage changes in livestock in TLUs vary
 501 significantly by bin ($p = 0.042$; rank transformed), with the middle bin higher than the
 502 bottom bin ($p = 0.1$) and the high bin higher than the bottom bin ($p = 0.06$). Annual
 503 percentage changes in fertilizer N ($p = 0.58$) and crop production ($p = 0.62$; rank
 504 transformed) did not vary by bin.

505



506
 507 **Figure 8.** Annual percentage changes in national mean annual NH₃ VCDs, burned area, CO
 508 VCDs, livestock, crop yield, and fertilizer N use for African countries with low, medium, or
 509 high rates of NH₃ VCD change. Error bars represent the standard error of the mean. See
 510 Table S1 for the list of countries in each bin and Fig. S19 for an expanded set of variables.

511
 512
 513 Instead of a direct agricultural relationship with changes in NH₃ VCDs, there is the
 514 possibility that changes in biomass burning are associated with changes in NH₃ VCDs.
 515 Although the differences in the annual percentage change in burned area were not
 516 significant among bins ($p=0.54$; rank transformed), the overall pattern is consistent with
 517 earlier results finding that a reduction in burned area across the northern biomass burning
 518 region was associated in part with the expansion of agriculture and presumed
 519 anthropogenic suppression of fire (Andela et al., 2017; Andela and van der Werf, 2014).
 520 However, burned area as measured by MODIS is likely an imperfect predictor for NH₃

521 emissions—as noted previously, MODIS underestimates burned area by a factor of 3 to 6
522 during shoulder seasons (Roteta et al., 2019), which is when fires are expected to emit
523 more reduced species such as NH_3 (Zheng et al., 2018). In contrast to burned area, the
524 annual change in column densities of CO—which tends to be co-emitted with NH_3 from
525 fires—differed significantly among bins ($p < 0.001$; rank transformed) and was significantly
526 higher in the high bin than in the low or medium bins ($p < 0.001$, post-hoc tests). The higher
527 annual CO changes in the high bin could be related to larger anthropogenic fossil fuel
528 emissions, but we see no difference among bins in growth rates of CO_2 emissions ($p = 0.48$;
529 Figure S19); such a difference would be expected if differences in economic development
530 were responsible for the CO differences. These results leave open the possibility that
531 changes in either biofuel emissions or biomass burning emissions—perhaps from smaller
532 fires not observed in the MODIS burned area product—may be primarily responsible for
533 the difference in CO between bins, and may be contributing to the differences in NH_3
534 between bins. Changes in NO_2 VCDs and SO_2 concentrations can affect the lifetime of NH_3
535 (the latter by changing SO_4 concentrations), but do not appear to make an important
536 contribution to the observed trends in NH_3 VCDs among bins (Fig. S19, SI text).
537 Temperature, likewise, does not appear to play an important role (SI text).

538

539 **4. Conclusion**

540 Using IASI, we have observed both increases and decreases in atmospheric NH_3
541 VCDs in different regions in Africa between 2008 and 2018, with different factors affecting
542 trends in different regions.

543 We observed increases in NH_3 VCDs in West Africa, which earlier work had
544 concluded was likely related to increased fertilizer use. Fertilizer is not typically applied in
545 West Africa until the start of the growing season—often April—but we find that much of
546 the NH_3 increase occurs during February and March, suggesting that increasing fertilizer
547 use is unlikely to provide a complete explanation for the NH_3 trend. Agriculture may
548 nevertheless play a role, with enhanced burned area and especially CO concentrations in
549 February suggestive of increased burning of crop stubble in preparation for planting during
550 this time of year. Fires in this region tend to emit a greater proportion of less oxidized

551 species such as NH_3 at the end of the dry season, consistent with a biomass burning source
552 for the increasing NH_3 VCDs.

553 Decreases in NH_3 VCDs were largest in South Sudan, especially over the Sudd
554 wetland, where NH_3 VCDs vary seasonally with the extent of area flooded. As the
555 temporarily flooded areas of the Sudd dry out each year, NH_3 VCDs increase as reduction in
556 soil moisture drives increased production and volatilization of NH_3 . The area of the Sudd
557 that is flooded each year varies, and from 2008 until 2015, the area that remains flooded
558 during the dry season generally increased, producing a positive overall trend for the period
559 of 2008 through 2017. This increase in the dry season flooded area drove a decrease in
560 NH_3 VCDs: with less soil drying out, the seasonal maxima in NH_3 VCDs were lower.
561 Although it is possible that conflict in South Sudan could contribute to changes in NH_3
562 VCDs, the timing and distribution of conflict events and human displacement suggest that
563 other factors are likely more important.

564 Modest increases in NH_3 VCDs were observed in the Lake Victoria region. This
565 region has experienced increases in agricultural area during the IASI observation period,
566 and these changes explained a large proportion of the variation in NH_3 VCDs across large
567 patches of the region, where biomass burning could not. We expect that both expansion
568 and intensification of agriculture in this region could contribute to the positive NH_3 trend.

569 Considering national-scale statistics, comparisons between equally sized bins of 17
570 countries each suggested that changes in biomass burning emissions and livestock
571 emissions could contribute to differences in NH_3 VCDs among countries, but variables
572 related to cropped agriculture such as cropped area or fertilizer N use did not appear to be
573 important factors at this scale. This may be because although fertilizer use has been
574 increasing in sub-Saharan Africa, it remains extremely low relative to other continents, and
575 relative to the levels needed to attain food security. Average fertilizer use in most
576 countries in the region is under $20 \text{ kg N ha}^{-1} \text{ yr}^{-1}$, and sometimes less than $5 \text{ kg N ha}^{-1} \text{ yr}^{-1}$.
577 Although recommended fertilizer rates are lower in most African countries than in the U.S.
578 or Europe, increasing N inputs to 50 or $100 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ would represent a major
579 perturbation to the regional N cycle, and potentially a large new source of NH_3 to the
580 atmosphere. West Africa is already a global NH_3 hotspot (Van Damme et al., 2018),
581 suggesting that encouraging policies that can help to limit NH_3 emissions during the early

582 stages of agricultural intensification in Africa may help mitigate potential impacts on the
583 atmosphere. Fortunately, agricultural practices such as sub-surface application of
584 fertilizer, which is already being promoted to smallholder farmers, can serve to both limit
585 NH₃ emissions also help to increase crop yields.

586 These past and anticipated future trends also make the case for expanding capacity
587 for atmospheric monitoring in sub-Saharan Africa. Although long-term monitoring
588 networks have been established in West Africa (Adon et al., 2010; Ossouhou et al., 2019) and
589 South Africa (Conradie et al., 2016) as part of the INDAAF network, it is mainly focused on
590 deposition and the spatio-temporal resolution of surface measurements is very coarse
591 when compared to the data available in other parts of the world, which will limit our ability
592 to understand how agricultural and socio-economic development in Africa affect the
593 atmosphere. Satellite observations can help to bridge some of these data gaps, but have
594 their own spatio-temporal limitations, and would further benefit from additional high-
595 quality surface observations for evaluation of retrieval products.

596

597 **Data availability:** All data used in this study are available from public sources, with the
598 exception of Sudd wetland extent, which is available by request from Courtney Di Vittorio. The
599 IASI NH₃ and CO data are available from The IASI <https://iasi.aeris-data.fr>. The NOAA Global
600 Surface Temperature Dataset is available at [https://data.nodc.noaa.gov/cgi-](https://data.nodc.noaa.gov/cgi-bin/iso?id=gov.noaa.ncdc:C01585)
601 [bin/iso?id=gov.noaa.ncdc:C01585](https://data.nodc.noaa.gov/cgi-bin/iso?id=gov.noaa.ncdc:C01585). MODIS burned area data are available from
602 <https://www.globalfireshdata.org/data.html>. MODIS agricultural area are available at
603 <https://lpdaac.usgs.gov/products/mcd12c1v006/>. TRMM 3B42 precipitation data are available
604 from <https://pmm.nasa.gov/data-access/downloads/trmm>. The Gridded Livestock of the World
605 data are available from <https://livestock.geo-wiki.org/home-2/>. Population density data for 2017
606 are available at <https://landscan.ornl.gov/downloads/2017>. FAO national crop production and
607 fertilizer N data are available at <http://www.fao.org/faostat/en/>. Data on conflict events from
608 ACLED are available at <https://acleddata.com/#/dashboard>. World Bank national statistics
609 on refugees and internally displaced people are available at <https://data.worldbank.org>.

610

611 **Author Contribution:** J.E.H. designed the study, conducted the analysis, and wrote the paper.
612 NA, ED, CD, MO, CG-L, KT, and SEB contributed to study design and edited the paper. LC, P-
613 FC, and MVD developed the original IASI trace gas retrievals and edited the paper.

614 The authors declare that they have no conflict of interest.

615
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621 **References**

622 van der A, R. J., Eskes, H. J., Boersma, K. F., van Noije, T. P. C., Van Roozendael, M., De Smedt,
623 I., Peters, D. H. M. U. and Meijer, E. W.: Trends, seasonal variability and dominant NO_x
624 source derived from a ten year record of NO₂ measured from space, *J. Geophys. Res.*
625 *Atmos.*, 113(4), D04302, doi:10.1029/2007JD009021, 2008.

626 Adon, M., Galy-Lacaux, C., Yoboué, V., Delon, C., Lacaux, J. P., Castera, P., Gardrat, E., Pienaar,
627 J., Al Ourabi, H., Laouali, D., Diop, B., Sigha-Nkamdjou, L., Akpo, a., Tathy, J. P., Lavenu, F. and
628 Mougin, E.: Long term measurements of sulfur dioxide, nitrogen dioxide, ammonia, nitric
629 acid and ozone in Africa using passive samplers, *Atmos. Chem. Phys.*, 10(15), 7467–7487,
630 doi:10.5194/acp-10-7467-2010, 2010.

631 AGRA: AGRA in 2008: Building on the New Momentum in African Agriculture., 2009.

632 Andela, N. and van der Werf, G. R.: Recent trends in African fires driven by cropland
633 expansion and El Niño to la Niña transition, *Nat. Clim. Chang.*, 4(9), 791–795,
634 doi:10.1038/nclimate2313, 2014.

635 Andela, N., Morton, D. C., Giglio, L., Chen, Y., Van Der Werf, G. R., Kasibhatla, P. S., DeFries, R.
636 S., Collatz, G. J., Hantson, S., Kloster, S., Bachelet, D., Forrest, M., Lasslop, G., Li, F., Mangeon,
637 S., Melton, J. R., Yue, C. and Randerson, J. T.: A human-driven decline in global burned area,
638 *Science (80-.)*, 356(6345), 1356–1362, doi:10.1126/science.aal4108, 2017.

639 Anker, H. T., Baaner, L., Backes, C., Keessen, A. and Möckel, S.: Comparison of ammonia
640 regulation in Germany , the Netherlands and Denmark – legal framework, Copenhagen.,

641 2018.

642 Bauer, S. E., Tsigaridis, K. and Miller, R.: Significant atmospheric aerosol pollution caused by
643 world food cultivation, *Geophys. Res. Lett.*, 43(10), 5394–5400,
644 doi:10.1002/2016GL068354, 2016.

645 Behera, S. N., Sharma, M., Aneja, V. P. and Balasubramanian, R.: Ammonia in the
646 atmosphere: A review on emission sources, atmospheric chemistry and deposition on
647 terrestrial bodies, *Environ. Sci. Pollut. Res.*, 20(11), 8092–8131, doi:10.1007/s11356-013-
648 2051-9, 2013.

649 Bouwman, A. F., Lee, D. S., Asman, W. A. H., Dentener, F. J., Van Der Hoek, K. W. and Olivier, J.
650 G. J.: A global high-resolution emission inventory for ammonia, *Global Biogeochem. Cycles*,
651 11(4), 561–587, 1997.

652 Bustamante, M. M. C., Medina, E., Asner, G. P., Nardoto, G. B. and Garcia-Montiel, D. C.:
653 Nitrogen cycling in tropical and temperate savannas, *Biogeochemistry*, 79(1–2), 209–237,
654 doi:10.1007/s10533-006-9006-x, 2006.

655 Cahoon, D. R., Stocks, B. J., Levine, J. S., Cofer, W. R. and O'Neill, K. P.: Seasonal distribution of
656 African savanna fires, *Nature*, 359(6398), 812–815, doi:10.1038/359812a0, 1992.

657 Chen, L.-W. A., Verburg, P., Shackelford, A., Zhu, D., Susfalk, R., Chow, J. C. and Watson, J. G.:
658 Moisture effects on carbon and nitrogen emission from burning of wildland biomass,
659 *Atmos. Chem. Phys.*, 10(14), 6617–6625, doi:10.5194/acp-10-6617-2010, 2010.

660 Chilonda, P. and Otte, J.: Indicators to monitor trends in livestock production at national,
661 regional and international levels, *Livest. Res. Rural Dev.*, 18(8), 1–12, 2006.

662 Clarisse, L., Clerbaux, C., Dentener, F., Hurtmans, D. and Coheur, P. F.: Global ammonia
663 distribution derived from infrared satellite observations, *Nat. Geosci.*, 2(7), 479–483,
664 doi:10.1038/ngeo551, 2009.

665 Clarisse, L., Van Damme, M., Gardner, W., Coheur, P.-F., Clerbaux, C., Whitburn, S., Hadji-
666 Lazaro, J. and Hurtmans, D.: Atmospheric ammonia (NH₃) emanations from Lake Natron's
667 saline mudflats, *Sci. Rep.*, 9, 4441, doi:10.1038/s41598-019-39935-3, 2019.

668 Cobo, J. G., Dercon, G. and Cadisch, G.: Nutrient balances in African land use systems across
669 different spatial scales: A review of approaches, challenges and progress, *Agric. Ecosyst.*
670 *Environ.*, 136(1–2), 1–15, doi:10.1016/j.agee.2009.11.006, 2010.

671 Conradie, E. H., Van Zyl, P. G., Pienaar, J. J., Beukes, J. P., Galy-Lacaux, C., Venter, A. D. and

672 Mkhathshwa, G. V.: The chemical composition and fluxes of atmospheric wet deposition at
673 four sites in South Africa, *Atmos. Environ.*, 146, 113–131,
674 doi:10.1016/j.atmosenv.2016.07.033, 2016.

675 Van Damme, M., Wichink Kruit, R. J., Schaap, M., Clarisse, L., Clerbaux, C., Coheur, P. F.,
676 Dammers, E., Dolman, A. J. and Erisman, J. W.: Evaluating 4 years of atmospheric ammonia
677 (NH₃) over Europe using IASI satellite observations and LOTOS-EUROS model results, *J.*
678 *Geophys. Res.*, 119(15), 9549–9566, doi:10.1002/2014JD021911, 2014a.

679 Van Damme, M., Clarisse, L., Heald, C. L., Hurtmans, D., Ngadi, Y., Clerbaux, C., Dolman, A. J.,
680 Erisman, J. W. and Coheur, P. F.: Global distributions, time series and error characterization
681 of atmospheric ammonia NH₃ from IASI satellite observations, *Atmos. Chem. Phys.*, 14(6),
682 2905–2922, doi:10.5194/acp-14-2905-2014, 2014b.

683 Van Damme, M., Clarisse, L., Dammers, E., Liu, X., Nowak, J. B., Clerbaux, C., Flechard, C. R.,
684 Galy-Lacaux, C., Xu, W., Neuman, J. A., Tang, Y. S., Sutton, M. A., Erisman, J. W. and Coheur, P.
685 F.: Towards validation of ammonia (NH₃) measurements from the IASI satellite, *Atmos.*
686 *Meas. Tech.*, 8(3), 1575–1591, doi:10.5194/amt-8-1575-2015, 2015.

687 Van Damme, M., Whitburn, S., Clarisse, L., Clerbaux, C., Hurtmans, D. and Coheur, P. F.:
688 Version 2 of the IASI NH₃ neural network retrieval algorithm: Near-real-time and
689 reanalysed datasets, *Atmos. Meas. Tech.*, 10(12), 4905–4914, doi:10.5194/amt-10-4905-
690 2017, 2017.

691 Van Damme, M., Clarisse, L., Whitburn, S., Hadji-Lazaro, J., Hurtmans, D., Clerbaux, C. and
692 Coheur, P.: Industrial and agricultural ammonia point sources exposed, *Nature*, 564, 99–
693 103, 2018.

694 Van Damme, M., Clarisse, L., Franco, B., Sutton, M. A., Erisman, J. W., Wichink Kruit, R., van
695 Zanten, M., Whitburn, S., Hadji-Lazaro, J., Hurtmans, D., Clerbaux, C. and Coheur, P.-F.:
696 Global, regional and national trends of atmospheric ammonia derived from a decadal
697 (2008-2018) satellite record, *Environ. Res. Lett.*, 2021.

698 Dammers, E., Palm, M., Van Damme, M., Vigouroux, C., Smale, D., Conway, S., Toon, G. C.,
699 Jones, N., Nussbaumer, E., Warneke, T., Petri, C., Clarisse, L., Clerbaux, C., Hermans, C.,
700 Lutsch, E., Strong, K., Hannigan, J. W., Nakajima, H., Morino, I., Herrera, B., Stremme, W.,
701 Grutter, M., Schaap, M., Kruit, R. J. W., Notholt, J., Coheur, P. F. and Erisman, J. W.: An
702 evaluation of IASI-NH₃ with ground-based Fourier transform infrared spectroscopy

703 measurements, *Atmos. Chem. Phys.*, 16(16), 10351–10368, doi:10.5194/acp-16-10351-
704 2016, 2016.

705 Denier Van Der Gon, H. and Bleeker, A.: Indirect N₂O emission due to atmospheric N
706 deposition for the Netherlands, *Atmos. Environ.*, 39(32), 5827–5838,
707 doi:10.1016/j.atmosenv.2005.06.019, 2005.

708 Dobson, J. E., Bright, E. A., Coleman, P. R., Durfee, R. C. and Worley, B. A.: A global population
709 database for estimating populations at risk, *Photogramm. Eng. Remote Sens.*, 66(7), 849–
710 857 [online] Available from: In Book..., 2000.

711 Doumbia, E. H. T., Liousse, C., Keita, S., Granier, L., Granier, C., Elvidge, C. D., Elguindi, N. and
712 Law, K.: Flaring emissions in Africa: Distribution, evolution and comparison with current
713 inventories, *Atmos. Environ.*, 199(November 2018), 423–434,
714 doi:10.1016/j.atmosenv.2018.11.006, 2019.

715 Elrys, A. S., Abdel-Fattah, M. K., Raza, S., Chen, Z. and Zhou, J.: Spatial trends in the nitrogen
716 budget of the African agro-food system over the past five decades, *Environ. Res. Lett.*,
717 14(12), doi:10.1088/1748-9326/ab5d9e, 2019.

718 European Commission Joint Research Centre (JRC)/Netherlands Environmental
719 Assessment Agency (PBL): Emission Database for Global Atmospheric Research (EDGAR),
720 release version 4.3.1, 2016.

721 Fan, Y. and van den Dool, H.: A global monthly land surface air temperature analysis for
722 1948-present, *J. Geophys. Res. Atmos.*, 113(1), D01103, doi:10.1029/2007JD008470, 2008.

723 FAO: FAO Statistics Database, FAOSTAT Stat. Database [online] Available from:
724 <http://www.fao.org/faostat/en/> (Accessed 1 January 2019), 2020.

725 Friedl, M. A., McIver, D. K., Hodges, J. C. F., Zhang, X. Y., Muchoney, D., Strahler, A. H.,
726 Woodcock, C. E., Gopal, S., Schneider, A., Cooper, A., Baccini, A., Gao, F. and Schaaf, C.: Global
727 land cover mapping from MODIS: Algorithms and early results, *Remote Sens. Environ.*,
728 83(1–2), 287–302, doi:10.1016/S0034-4257(02)00078-0, 2002.

729 Galy-Lacaux, C. and Delon, C.: Nitrogen emission and deposition budget in West and Central
730 Africa, *Environ. Res. Lett.*, 9(12), doi:10.1088/1748-9326/9/12/125002, 2014.

731 Gbadegesin, A. and Olusesi, B. B.: Effects of land clearing methods on soil physical and
732 hydrological properties in southwestern Nigeria, *Environmentalist*, 14(4), 297–303
733 [online] Available from: <http://www.scopus.com/inward/record.url?eid=2-s2.0->

734 0022858519&partnerID=40&md5=3051b1dd91ff2990e37fe7466872923e, 1994.
735 George, M., Clerbaux, C., Hurtmans, D., Turquety, S., Coheur, P. F., Pommier, M., Hadji-
736 Lazaro, J., Edwards, D. P., Worden, H., Luo, M., Rinsland, C. and McMillan, W.: Carbon
737 monoxide distributions from the IASI/METOP mission: Evaluation with other space-borne
738 remote sensors, *Atmos. Chem. Phys.*, 9(21), 8317–8330, doi:10.5194/acp-9-8317-2009,
739 2009.

740 Giglio, L., Boschetti, L., Roy, D. P., Humber, M. L. and Justice, C. O.: The Collection 6 MODIS
741 burned area mapping algorithm and product, *Remote Sens. Environ.*, 217, 72–85,
742 doi:10.1016/j.rse.2018.08.005, 2018.

743 Global Internal Displacement Monitoring Centre: Global Internal Displacement Database,
744 [online] Available from: [https://www.internal-displacement.org/database/displacement-](https://www.internal-displacement.org/database/displacement-data)
745 [data](https://www.internal-displacement.org/database/displacement-data), 2020.

746 Goode, J. G., Yokelson, R. J., Susott, R. A. and Ward, D. E.: Trace gas emissions from
747 laboratory biomass fires measured by open-path Fourier transform infrared spectroscopy,
748 *J. Chem. Inf. Model.*, 104(D17), 21237–21245, 1999.

749 Guo, X., Clarisse, L., Wang, R., Van Damme, M., Whitburn, S., Coheur, P., Clerbaux, C., Franco,
750 B., Pan, D., Golston, L. M., Wendt, L., Sun, K., Tao, L., Miller, D., Mikoviny, T., Müller, M.,
751 Wisthaler, A., Tevlin, A. G., Murphy, J. G., Nowak, J. B., Roscioli, J. R., Volkamer, R., Kille, N.,
752 Neuman, J. A., Eilerman, S. J., Crawford, J. H., Yacovitch, T. I., Barrick, J. D., Scarino, A. J. and
753 Zondlo, M. A.: Validation of IASI satellite ammonia observations at the pixel scale using in-
754 situ vertical profiles, *J. Geophys. Res. Atmos.*, 126, e2020JD033475,
755 doi:10.1029/2020jd033475, 2021.

756 Hazell, P. and Wood, S.: Drivers of change in global agriculture., *Philos. Trans. R. Soc. Lond.*
757 *B. Biol. Sci.*, 363(1491), 495–515, doi:10.1098/rstb.2007.2166, 2008.

758 Hickman, J. E., Dammers, E., Galy-Lacaux, C. and Van Der Werf, G. R.: Satellite evidence of
759 substantial rain-induced soil emissions of ammonia across the Sahel, *Atmos. Chem. Phys.*,
760 18(22), 16713–16727, doi:10.5194/acp-18-16713-2018, 2018.

761 Hickman, J. E., Andela, N., Tsigaridis, K., Galy-Lacaux, C., Osohou, M. and Bauer, S. E.:
762 Reductions in NO₂ burden over north equatorial Africa from decline in biomass burning in
763 spite of growing fossil fuel use, 2005 to 2017, *Proc. Natl. Acad. Sci.*, 118(7), e2002579118,
764 doi:10.1073/pnas.2002579118, 2021a.

765 Hickman, J. E., Dammers, E., Galy-Lacaux, C., Ossohou, M. and Bauer, S. E.: Continental and
766 ecoregion-specific drivers of atmospheric NO₂ and NH₃ seasonality over Africa revealed by
767 satellite observations, *Global Biogeochem. Cycles*, 35, e2020GB006916,
768 doi:10.1029/2020GB006916, 2021b.

769 Huffman, G. J., Adler, R. F., Bolvin, D. T., Gu, G., Nelkin, E. J., Bowman, K. P., Hong, Y., Stocker,
770 E. F. and Wolff, D. B.: The TRMM Multisatellite Precipitation Analysis (TMPA): Quasi-Global,
771 Multiyear, Combined-Sensor Precipitation Estimates at Fine Scales, *J. Hydrometeorol.*, 8(1),
772 38–55, doi:10.1175/jhm560.1, 2007.

773 Hurtmans, D., Coheur, P. F., Wespes, C., Clarisse, L., Scharf, O., Clerbaux, C., Hadji-Lazaro, J.,
774 George, M. and Turquety, S.: FORLI radiative transfer and retrieval code for IASI, *J. Quant.*
775 *Spectrosc. Radiat. Transf.*, 113(11), 1391–1408, doi:10.1016/j.jqsrt.2012.02.036, 2012.

776 Idris, I.: Livestock and conflict in South Sudan - K4D Helpdesk Report 484, Brighton.
777 [online] Available from:
778 https://assets.publishing.service.gov.uk/media/5c6abdec40f0b61a22792fd5/484_Livestock_and_Conflict_in_South_Sudan.pdf, 2018.

780 Kanter, D. R.: Nitrogen pollution: a key building block for addressing climate change, *Clim.*
781 *Change*, 147(1–2), 11–21, doi:10.1007/s10584-017-2126-6, 2018.

782 Kerzenmacher, T., Dils, B., Kumps, N., Blumenstock, T., Clerbaux, C., Coheur, P. F., Demoulin,
783 P., García, O., George, M., Griffith, D. W. T., Hase, F., Hadji-Lazaro, J., Hurtmans, D., Jones, N.,
784 Mahieu, E., Notholt, J., Paton-Walsh, C., Raffalski, U., Ridder, T., Schneider, M., Servais, C. and
785 De Mazière, M.: Validation of IASI FORLI carbon monoxide retrievals using FTIR data from
786 NDACC, *Atmos. Meas. Tech.*, 5(11), 2751–2761, doi:10.5194/amt-5-2751-2012, 2012.

787 Korontzi, S., McCarty, J., Loboda, T., Kumar, S. and Justice, C.: Global distribution of
788 agricultural fires in croplands from 3 years of Moderate Resolution Imaging
789 Spectroradiometer (MODIS) data, *Global Biogeochem. Cycles*, 20(2), GB2021,
790 doi:10.1029/2005GB002529, 2006.

791 Krupa, S. V.: Effects of atmospheric ammonia (NH₃) on terrestrial vegetation: A review,
792 *Environ. Pollut.*, 124(2), 179–221, doi:10.1016/S0269-7491(02)00434-7, 2003.

793 Lee, C., Martin, R. V., Van Donkelaar, A., Lee, H., Dickerson, R. R., Hains, J. C., Krotkov, N.,
794 Richter, A., Vinnikov, K. and Schwab, J. J.: SO₂ emissions and lifetimes: Estimates from
795 inverse modeling using in situ and global, space-based (SCIAMACHY and OMI)

796 observations, *J. Geophys. Res. Atmos.*, 116(6), 1–13, doi:10.1029/2010JD014758, 2011.

797 Lelieveld, J., Evans, J. S., Fnais, M., Giannadaki, D. and Pozzer, A.: The contribution of outdoor
798 air pollution sources to premature mortality on a global scale, *Nature*, 525(7569), 367–371,
799 doi:10.1038/nature15371, 2015.

800 Masso, C., Nziguheba, G., Mtutegi, J., Galy-Lacaux, C., Wendt, K., Butterbach-Bahl, K., Wairegi,
801 L. and Datta, A.: Soil fertility Management in Sub-Saharan Africa, in *Sustainable Agriculture*
802 *Reviews*, edited by E. Lichtfouse, p. 304, Springer International Publishing., 2017.

803 Nicholson, S., Some, B., McCollum, J., Nelkin, E., Klotter, D., Berte, Y., Diallo, B., Gaye, I.,
804 Kpabeba, G., Ndiaye, O., Noukpozoukou, J., Tanu, M., Thiam, A., Toure, A. and Traore, A.:
805 Validation of TRMM and other rainfall estimates with a high-density gauge dataset for West
806 Africa. Part II: Validation of TRMM rainfall products, *J. Appl. Meteorol.*, 42, 1355–1368,
807 doi:Article, 2003.

808 Osohou, M., Galy-Lacaux, C., Yoboué, V., Hickman, J. E., Gardrat, E., Adon, M., Darras, S.,
809 Laouali, D., Akpo, A., Ouafou, M., Diop, B. and Opepa, C.: Trends and seasonal variability of
810 atmospheric NO₂ and HNO₃ concentrations across three major African biomes inferred
811 from long-term series of ground-based and satellite measurements, *Atmos. Environ.*, 207,
812 148–166, 2019.

813 Pamela A. Matson, McDowell, W. H., Townsend, A. R. and Vitousek, P. M.: The globalization
814 of N deposition: ecosystem consequences in tropical environments, *Biogeochemistry*, 46,
815 67–83 [online] Available from: [http://academic.engr.arizona.edu/HWR/Brooks/GC572-](http://academic.engr.arizona.edu/HWR/Brooks/GC572-2004/readings/matson.pdf)
816 [2004/readings/matson.pdf](http://academic.engr.arizona.edu/HWR/Brooks/GC572-2004/readings/matson.pdf), 1999.

817 Pommier, M., Law, K. S., Clerbaux, C., Turquety, S., Hurtmans, D., Hadji-Lazaro, J., Coheur, P.
818 F., Schlager, H., Ancellet, G., Paris, J. D., Nédélec, P., Diskin, G. S., Podolske, J. R., Holloway, J. S.
819 and Bernath, P.: IASI carbon monoxide validation over the Arctic during POLARCAT spring
820 and summer campaigns, *Atmos. Chem. Phys.*, 10(21), 10655–10678, doi:10.5194/acp-10-
821 10655-2010, 2010.

822 Pope, A., Burnett, R., Thun, M., EE, C., D, K., I, K. and GD, T.: Long-term Exposure to Fine
823 Particulate Air Pollution, *J. Am. Med. Assoc.*, 287(9), 1132–1141,
824 doi:10.1001/jama.287.9.1132, 2002.

825 Raleigh, C., Linke, A., Hegre, H. and Karlsen, J.: Introducing ACLED: An Armed Conflict
826 Location and Event Dataset: Special Data Feature, *J. Peace Res.*, 47(5), 651–660,

827 doi:<https://doi.org/10.1177/0022343310378914>, 2010.

828 Ramo, R., Roteta, E., Bistinas, I., van Wees, D., Bastarrika, A., Chuvieco, E. and van der Werf,
829 G. R.: African burned area and fire carbon emissions are strongly impacted by small fires
830 undetected by coarse resolution satellite data, *Proc. Natl. Acad. Sci. U. S. A.*, 118(9), 1–7,
831 doi:10.1073/pnas.2011160118, 2021.

832 Robinson, T. P., Wint, G. R. W., Conchedda, G., Van Boeckel, T. P., Ercoli, V., Palamara, E.,
833 Cinardi, G., D’Aietti, L., Hay, S. I. and Gilbert, M.: Mapping the Global Distribution of
834 Livestock, *PLoS One*, 9(5), e96084, doi:10.1371/journal.pone.0096084, 2014.

835 Roteta, E., Bastarrika, A., Padilla, M., Storm, T. and Chuvieco, E.: Development of a Sentinel-2
836 burned area algorithm: Generation of a small fire database for sub-Saharan Africa, *Remote
837 Sens. Environ.*, 222(November 2018), 1–17, doi:10.1016/j.rse.2018.12.011, 2019.

838 Rufino, M. C., Rowe, E. C., Delve, R. J. and Giller, K. E.: Nitrogen cycling efficiencies through
839 resource-poor African crop-livestock systems, *Agric. Ecosyst. Environ.*, 112(4), 261–282,
840 doi:10.1016/j.agee.2005.08.028, 2006.

841 Sauvage, B., Gheusi, F., Thouret, V., Cammas, J. P., Duron, J., Escobar, J., Mari, C., Mascart, P.
842 and Pont, V.: Medium-range mid-tropospheric transport of ozone and precursors over
843 Africa: Two numerical case studies in dry and wet seasons, *Atmos. Chem. Phys.*, 7(20),
844 5357–5370, doi:10.5194/acp-7-5357-2007, 2007.

845 Stevens, C. J., David, T. I. and Storkey, J.: Atmospheric nitrogen deposition in terrestrial
846 ecosystems: Its impact on plant communities and consequences across trophic levels,
847 *Funct. Ecol.*, 32(7), 1757–1769, doi:10.1111/1365-2435.13063, 2018.

848 Sutton, M. A., Reis, S., Riddick, S. N., Dragosits, U., Nemitz, E., Theobald, M. R., Tang, Y. S.,
849 Braban, C. F., Vieno, M., Dore, A. J., Mitchell, R. F., Wanless, S., Daunt, F., Fowler, D., Blackall,
850 T. D., Milford, C., Flechard, C. R., Loubet, B., Massad, R., Cellier, P., Personne, E., Coheur, P. F.,
851 Clarisse, L., Van Damme, M., Ngadi, Y., Clerbaux, C., Skøth, C. A., Geels, C., Hertel, O., Wichink
852 Kruit, R. J., Pinder, R. W., Bash, J. O., Walker, J. T., Simpson, D., Horváth, L., Misselbrook, T. H.,
853 Bleeker, A., Dentener, F. and de Vries, W.: Towards a climate-dependent paradigm of
854 ammonia emission and deposition, *Philos. Trans. R. Soc. B Biol. Sci.*, 368(1621), 20130166–
855 20130166, doi:10.1098/rstb.2013.0166, 2013.

856 Tian, D. and Niu, S.: A global analysis of soil acidification caused by nitrogen addition,
857 *Environ. Res. Lett.*, 10(2), 024019, doi:10.1088/1748-9326/10/2/024019, 2015.

858 USDA Agricultural Air Quality Task Force: Ammonia Emissions : What To Know Before You
859 Regulate, Washington, DC. [online] Available from:
860 <http://www.nrcs.usda.gov/wps/portal/nrcs/detail/national/air/taskforce/?cid=stelprdb1>
861 268645, 2014.

862 Vitousek, P., Naylor, R., Crews, T., David, M., Drinkwater, L., Holland, E., Johnes, P.,
863 Katzenberger, J., Martinelli, L. A., Matson, P. A., Nziguheba, G., Ojima, D., Palm, C. A.,
864 Robertson, G., Sanchez, P., Townsend, A. and Zhang, F.: Nutrient Imbalances in Agricultural
865 Development, *Science* (80-.), 324, 1519–1520, 2009.

866 Di Vittorio, C. A. and Georgakakos, A. P.: Land cover classification and wetland inundation
867 mapping using MODIS, *Remote Sens. Environ.*, 204(May 2017), 1–17,
868 doi:10.1016/j.rse.2017.11.001, 2018.

869 Vrieling, A., de Beurs, K. M. and Brown, M. E.: Variability of African farming systems from
870 phenological analysis of NDVI time series, *Clim. Change*, 109(3–4), 455–477,
871 doi:10.1007/s10584-011-0049-1, 2011.

872 De Wachter, E., Barret, B., Le Flochmoën, E., Pavelin, E., Matricardi, M., Clerbaux, C., Hadji-
873 Lazaro, J., George, M., Hurtmans, D., Coheur, P. F., Nedelec, P. and Cammas, J. P.: Retrieval of
874 MetOp-A/IASI CO profiles and validation with MOZAIC data, *Atmos. Meas. Tech.*, 5(11),
875 2843–2857, doi:10.5194/amt-5-2843-2012, 2012.

876 Warner, J. X., Dickerson, R. R., Wei, Z., Strow, L. L., Wang, Y. and Liang, Q.: Increased
877 atmospheric ammonia over the world’s major agricultural areas detected from space,
878 *Geophys. Res. Lett.*, 44(6), 2875–2884, doi:10.1002/2016GL072305, 2017.

879 Whitburn, S., Van Damme, M., Kaiser, J. W., Van Der Werf, G. R., Turquety, S., Hurtmans, D.,
880 Clarisse, L., Clerbaux, C. and Coheur, P. F.: Ammonia emissions in tropical biomass burning
881 regions: Comparison between satellite-derived emissions and bottom-up fire inventories,
882 *Atmos. Environ.*, 121, 42–54, doi:10.1016/j.atmosenv.2015.03.015, 2015.

883 World Bank: World Bank Open Data, World Bank Open Data [online] Available from:
884 <https://www.data.worldbank.org> (Accessed 2 February 2019), n.d.

885 Yegbemey, R. N., Kabir, H., Awoye, O. H. R., Yabi, J. A. and Paraiso, A. A.: Managing the
886 agricultural calendar as coping mechanism to climate variability: A case study of maize
887 farming in northern Benin, West Africa, *Clim. Risk Manag.*, 3, 13–23,
888 doi:10.1016/j.crm.2014.04.001, 2014.

889 Yokelson, R. J., Christian, T. J., Karl, T. G. and Guenther, A.: The tropical forest and fire
890 emissions experiment: laboratory fire measurements and synthesis of campaign, *Rev. Int.*
891 *Acupunt.*, 8, 3509–3527, doi:10.1016/s1887-8369(09)71579-0, 2008.

892 Zheng, B., Chevallier, F., Ciais, P., Yin, Y. and Wang, Y.: On the Role of the Flaming to
893 Smoldering Transition in the Seasonal Cycle of African Fire Emissions, *Geophys. Res. Lett.*,
894 45(21), 11,998-12,007, doi:10.1029/2018GL079092, 2018.

895