| 1 | Changes in biomass burning, wetland extent, or agriculture |
|------------------|--|
| 2 | drive atmospheric NH3 trends in several African regions |
| 3 | |
| 4 | |
| 5 | |
| 6 7 8 9 | Jonathan E. Hickman ^{1*,} Niels Andela ^{2†} , Enrico Dammers ³ , Lieven Clarisse ⁴ , Pierre-François Coheur ⁴ , Martin Van Damme ⁴ , Courtney Di Vittorio ⁵ , Money Ossohou ⁶ , Corinne Galy- Lacaux ⁷ , Kostas Tsigirdis ^{1,8} , Susanne Bauer ¹ |
| 10 | ¹ NASA Goddard Institute for Space Studies, New York, USA |
| 11 | ² NASA Goddard Space Flight Center, Beltsville, USA |
| 12 | ³ Air Quality Research Division, Environment and Climate Change Canada, Toronto, |
| 13 | Canada |
| 14 | ⁴ Université libre de Bruxelles (ULB), Service de Chimie Quantique et Photophysique, |
| 15 | Atmospheric Spectroscopy, Brussels, Belgium |
| 16 | ⁵ Wake Forest University, Winston-Salem, USA |
| 17 | ⁶ Laboratoire des Sciences de la Matière, de l'Environnement et de l'Energie Solaire, |
| 18 | Université Félix Houphouët-Boigny, Abidjan, Côte d'Ivoire |
| 19 | ⁷ Laboratoire d'Aérologie, Université Toulouse III Paul Sabatier / CNRS, France |
| 20 | ⁸ Columbia University, New York, USA |
| 21 22 | [†] Now at School of Earth and Ocean Sciences, Cardiff University, Cardiff, UK |
| 23 | *Correspondence to: jonathan.e.hickman@nasa.gov |
| 24 | |
| 25 | Abstract |
| 26 | Atmospheric ammonia (NH_3) is a precursor to fine particulate matter and a source |
| 27 | of nitrogen (N) deposition that can adversely affect ecosystem health. The main sources of |
| 28 | NH_3 —agriculture and biomass burning—are undergoing or expected to undergo |
| 29 | substantial changes in Africa. Although evidence of increasing NH_3 over parts of Africa has |
| 30 | been observed, the mechanisms behind these trends are not well understood. Here we use |

31 observations of atmospheric NH₃ vertical column densities (VCDs) from the Infrared 32 Atmospheric Sounding Interferometer (IASI) along with other satellite observations of the 33 land surface and atmosphere to evaluate how NH₃ concentrations have changed over Africa 34 from 2008 through 2018, and what has caused those changes. In West Africa NH₃ VCDs are 35 observed to increase during the late dry season, with increases of over 7% yr⁻¹ in Nigeria 36 during February and March (p < 0.01). These positive trends are associated with increasing 37 burned area and CO trends during these months, likely related to agricultural preparation. 38 Increases are also observed in the Lake Victoria Basin, where they are associated with 39 expanding agricultural area. In contrast, NH₃ VCDs declined over the Sudd wetlands in 40 South Sudan by over 2% yr⁻¹ (p=0.20). Annual maxima in NH₃ VCDs in South Sudan occur 41 during February through May and are associated with drying of temporarily flooded 42 wetland soils, which favor emissions of NH₃. The change in mean NH₃ VCDs over the Sudd 43 is strongly correlated with variation in wetland extent in the Sudd: in years when more 44 area remained flooded during the dry season, NH_3 concentrations were higher (r=0.65, 45 p=0.04). Relationships between biomass burning and NH₃ may be observed when 46 evaluating national-scale statistics: countries with the highest rates of increasing NH₃ VCDs 47 also had high rates of growth in CO VCDs; burned area displayed a similar pattern, though 48 not significantly. Livestock numbers were also higher in countries with intermediate or 49 high rates of NH₃ VCD growth. Fertilizer use in Africa is currently low but growing; 50 implementing practices that can limit NH₃ losses from fertilizer as agriculture is intensified 51 may help mitigate impacts on health and ecosystems.

- 52 53
 - **1. Introduction:**
- 54

Ammonia (NH₃), a reactive nitrogen (N) trace gas, plays a number of important roles in the atmosphere, with implications for human health, climate, and ecosystems. Once in the atmosphere, NH₃ contributes to the production of inorganic aerosols, the primary constituents of fine particulate matter and a serious health hazard (Bauer et al., 2016; Lelieveld et al., 2015; Pope et al., 2002). NH₃ can also be deposited to downwind ecosystems, contributing to eutrophication, soil acidification, vegetation damage, productivity declines, reductions in biodiversity, and indirect greenhouse gas emissions 62 (Denier Van Der Gon and Bleeker, 2005; Krupa, 2003; Matson et al., 1999; Stevens et al.,

63 2018; Tian and Niu, 2015).

Although NH₃ is emitted from natural soils, agriculture is by far the largest source of
NH₃ globally (Behera et al., 2013; Bouwman et al., 1997). Urea fertilizer and livestock
excreta are particularly important substrates for NH₃ formation, and can be volatilized
quickly under favorable environmental conditions (Bouwman et al., 1997). In all soils, NH₃
is formed in solution following the dissociation of ammonium (NH₄⁺; Eq. 1).

- 69
- 70

 $NH_{4^{+}} + OH^{-} \leftrightarrow H_2O + NH_3$ (Eq. 1)

71

72 Soil NH₃ production is temperature-dependent, doubling with every 5°C temperature 73 increase, though the actual soil NH₃ flux is determined in part by plant and soil physiological 74 and physical factors (Sutton et al., 2013). On average, fertilizer use has been extremely low 75 in sub-Saharan Africa—often an order of magnitude or more lower than typical in Europe, 76 the United States, or China (Hazell and Wood, 2008; Vitousek et al., 2009). Livestock manure 77 N content also tends to be very low in sub-Saharan Africa (Rufino et al., 2006); the low 78 fertilizer use suggests that natural soils (as opposed to agricultural soils) may be a more 79 important source in the region than elsewhere in the world. However, agricultural 80 intensification and increasing fertilizer use has been a central policy focus for many African 81 countries, with national and regional efforts to increase N inputs by an order of magnitude 82 or more (AGRA, 2009).

83 After agriculture, biomass burning is the most important source of NH₃ globally 84 (Bouwman et al., 1997), with roughly 60 to 70% of global NH₃ emissions from fires occurring 85 in Africa (Cahoon et al., 1992; Whitburn et al., 2015). The amount of NH_3 emitted from 86 biomass fires is controlled primarily by the type of burning that occurs. N in fuel is present 87 predominantly in a chemically reduced state, and NH_3 is emitted in greater quantities from 88 low temperature, smoldering combustion in which fuel N is incompletely oxidized (Goode et 89 al., 1999; Yokelson et al., 2008). Fuel moisture content, which can help determine whether 90 combustion is smoldering or flaming, is thus an important determinant of biomass burning 91 NH₃ emissions (Chen et al., 2010).

92 In contrast to other reactive N gases such as NO_x (nitric oxide + nitrogen dioxide), 93 NH₃ emissions are typically unregulated outside of Europe (Anker et al., 2018; Kanter, 94 2018; USDA Agricultural Air Quality Task Force, 2014), and substantial increasing trends 95 have been observed by the NASA Atmospheric InfraRed Sounder (AIRS) and the Infrared 96 Atmospheric Sounding Interferometer (IASI) over many of the world's major agricultural 97 and biomass burning regions during the 21st century (Van Damme et al., 2021; Warner et 98 al., 2017). West Africa has been identified as an important NH_3 source region (Van Damme 99 et al., 2018), where a trend of increasing NH_3 concentrations over 2002 to 2013 has been 100 attributed at least in part to increased fertilizer use (Van Damme et al., 2021; Warner et al., 101 2017). Increasing trends have also been observed over central Africa, and attributed to 102 higher rates of biomass burning (Van Damme et al., 2021; Warner et al., 2017). However, 103 the studies by Warner et al. (2017) and Van Damme et al. (2021) were global in nature, 104 and as such could not include detailed explorations of the drivers of trends such as 105 consideration of emission seasonality or the geographic distribution of emission drivers, 106 which are particularly important across large parts of Africa where both biomass burning 107 and soils are potentially important sources (van der A et al., 2008).

Here we use a ten-year satellite record to evaluate trends in atmospheric NH₃
concentrations over Africa from 2008 through 2017, including detailed examination of
three regions where changes are pronounced: West Africa, the Lake Victoria Region, and
South Sudan.

112

113 **2.** Data and Methods

114 **2.1 Global gridded data**

115 Multiple data products were used, including satellite observations and spatial116 datasets:

-IASI-A, launched aboard the European Space Agency's MetOp-A in 2006, provides
measurements of atmospheric NH₃ and carbon monoxide (CO) twice a day (9:30 in the
morning and evening, Local Solar Time at the equator). Here we use morning observations,
when the thermal contrast is more favorable for retrievals (Clarisse et al., 2009; Van
Damme et al., 2014a). The NH₃ retrieval product used (ANNI-NH₃-v3R) follows a neural
network retrieval approach. We refer to Van Damme et al. (2017) and Van Damme et al.

123 (2021) for a detailed description of the algorithm. For CO, we used the product obtained 124 with the FORLI v20140922 retrieval algorithm (Hurtmans et al., 2012). Given the absence 125 of hourly or even daily observations of NH₃ concentrations in sub-Saharan Africa, the 126 detection limit of IASI is difficult to determine with certainty. However, the region 127 experiences high thermal contrast, and IASI seems to be able to reliably observe NH₃ down 128 to 1 to 2 ppb at the surface (Clarisse et al., 2009; Van Damme et al., 2014b). We gridded the 129 Level-2 IASI NH₃ and CO products to $0.25^{\circ} \times 0.25^{\circ}$ resolution to match the resolution of 130 other data used in the analysis. We used a conventional binning approach based on the 131 center of each satellite footprint. We did not apply an averaging weight. Quality control 132 procedures were followed as detailed in van Damme et al. 2017 and Van Damme et al., 133 2021. Specifically, the screening of retrievals included filtering of retrievals where cloud 134 cover is over 10%, where the total column density is below zero and the absolute value of 135 the hyperspectral range index (HRI) is above 1.5, and where the ratio of the total column 136 density to HRI is larger than 1.5×10^{16} molecules cm⁻².

137 The IASI products have been validated using ground-based Fourier transform infrared (FTIR) observations of NH₃ total columns, with robust correlations at sites with 138 139 high NH₃ concentrations, but lower at sites where atmospheric concentrations approach 140 IASI's detection limits (Dammers et al., 2016; Guo et al., 2021). Compared to the FTIR 141 observations, total columns from previous IASI NH₃ products (IASI-LUT and IASI-NNv1) 142 are biased low by $\sim 30\%$ which varies per region depending on the local concentrations. 143 Although FTIR observations are absent from Africa, earlier work has shown fair agreement 144 between previous versions of IASI total column densities and surface observations of NH₃ 145 using passive samplers across the International Network to study Deposition and 146 Atmospheric chemistry in AFrica (INDAAF) network in West Africa (Van Damme et al., 147 2015), including in observations of seasonal variation (Hickman et al., 2018; Ossohou et al., 148 2019). Validation of the IASI CO product using surface, aircraft, and satellite observations 149 have found total columns to have an error that is generally below 10-15% in the tropics 150 and mid-latitudes (George et al., 2009; Kerzenmacher et al., 2012; Pommier et al., 2010; De 151 Wachter et al., 2012). The IASI NH_3 and CO products were used for the years 2008—the 152 first full year of data available—to the end of 2018. Random errors in observations can be

assumed to cancel out in the annual mean, which is what we used in our analysis. With the

- assumption that random errors cancel out, only systematic errors related to tropospheric
- 155 vertical column contents remain; these systematic errors do not contribute to uncertainty
- 156 in trend analyses. In addition, we first take monthly averages based on all daily
- 157 observations within a given month before calculating seasonal means to minimize any
- 158 potential effects of temporal variability in cloud cover.
- -The Tropical Rainfall Measuring Mission (TRMM) daily precipitation product (3B42)
 is based on a combination of TRMM observations, geo-synchronous infrared observations,
 and rain gauge observations (Huffman et al., 2007). Independent rain gauge observations
 from West Africa have been used to validate the product, with no indication of bias in the
 product (Nicholson et al., 2003).
- NOAA Global Surface Temperature Dataset, a 0.5° gridded 2m monthly land surface
 temperature product (Fan and van den Dool, 2008). The data set is based on a combination
 of station observations from the Global Historical Climatology Network version 2 and the
 Climate Anomaly Monitoring System (GHCN_CAMS), and uses an anomaly interpolation
 approach which relies on observation-based reanalysis data to derive spatio-temporal
 variation in temperature lapse rates for topographic temperature adjustment.
- 500m MCD64A1 collection 6 Moderate Resolution Imaging Spectroradiometer
 (MODIS) burned area product for the period 2008-2018 (Giglio et al., 2018). The burned
 area data are aggregated by month and gridded to 0.25° resolution, and do not include
 burned area from small fires.
- MODIS MCD12C1 (collection 5) land cover product, which provides the percentage
 of cropped area in each 0.25° grid cell (Friedl et al., 2002). In Africa, agriculture is often
 practiced in complex mosaics of agricultural and natural land cover, so we used both the crop
 and crop/natural area mosaic MODIS classifications as agricultural area in our analysis.
- -We also used data on the spatio-temporal distribution of armed conflict events from
 the Armed Conflict Location & Event Data Project (ACLED; Raleigh et al., 2010). We included
 data for both violent and non-violent conflict events over the period 2008-2017.

181 **2.2 Sudd wetland extent**

182 Monthly flooded area extents of the Sudd Wetland, South Sudan from 2000 to 2017 183 were derived from 8-day composite MODIS land surface reflectance imagery (MOD09A1); 184 data from 2005 through 2017 were used in the analyses. We refer to Di Vittorio and 185 Georgakakos (2018) for a detailed description of the classification procedure designed to 186 retrieve these data. In summary, monthly flood maps were obtained through a two-stage 187 classification procedure. The first stage used the full 18-year data set to produce a wetland 188 land cover map that distinguishes between wetland vegetation classes and their flooding 189 regimes (permanently flooded, seasonally flooded, or non-flooded). The second stage compares seasonally flooded pixels from each vegetation class to their non-flooded 190 191 counterparts on a monthly basis to identify the timing and duration of flooding for each pixel. 192 These data were originally derived to calibrate a hydrologic model of the Sudd that is 193 dependent on Nile flows; therefore, a connectivity algorithm was applied to ensure that all 194 flooded pixels were physically connected to the Nile River. A few adjustments have been 195 made to the previously published dataset for the application of this study. The classification 196 algorithm has been improved to more accurately capture the inter-annual fluctuations in the 197 permanently flooded areas. The dataset was also extended through 2017, and the total 198 flooded area was quantified prior to applying the connectivity algorithm. The magnitudes of 199 the monthly flooded area estimates are now substantially larger because they include areas 200 flooded from local runoff in addition to areas flooded by the Nile River.

201

202

2.3 Spatial and national analyses

We evaluated spatial relationships between mean annual tropospheric NH₃ concentration and several independent variables at 0.25° resolution: population density, livestock density, and cropped area. Population density and livestock density data are not available as time series suitable for trend analysis, so we use single year values in our analyses. We calculated population density based on the 2017 version of the US Department of Energy's Gridded Landscan population dataset (Dobson et al., 2000;

209 available at https://landscan.ornl.gov). Livestock density was based on the FAO global 210 gridded livestock dataset for the year 2007 (Robinson et al., 2014). Before analysis, we 211 converted the livestock densities of chickens, goats, pigs, and sheep to tropical livestock 212 units (TLU), using values of 0.01, 0.1, 0.2, and 0.1 TLU, respectively; North African cattle 213 were converted using a factor of 0.7, whereas sub-Saharan cattle were converted using a 214 factor of 0.5 (Chilonda and Otte, 2006). For cropped area, we used the MODIS MCD12C1 215 (collection 5) land cover product as described above. We conducted spatial analyses by 216 establishing a map of 1.5° grid cells and calculating the correlation between the value of 217 each independent variable and NH₃ for all 0.25° grid cells within the larger grid cells (N = 218 144 including water grid cells, though these were excluded from the analysis).

219

220 National data on annual livestock numbers, crop production, and fertilizer N use were 221 obtained from the UN Food and Agriculture Organization FAOSTAT for 51 African countries 222 (FAO, 2020). Livestock data consisting of sheep, goats, cattle, and pigs were converted to 223 tropical livestock units as described above, and buffaloes were converted using a conversion 224 factors of 0.7 (Chilonda and Otte, 2006). National emissions of CO₂ were obtained from 225 World Bank Open Data (World Bank, 2019). National-level mean annual cropland area, 226 burned area, and atmospheric NH₃ and CO concentrations were also calculated for each of 227 the 51 countries from the spatial datasets described above. Countries were sorted into three 228 bins based on whether their relative change in mean annual NH_3 concentration was low, 229 medium or high, and means and standard errors were calculated for each of the three 17-230 country bins.

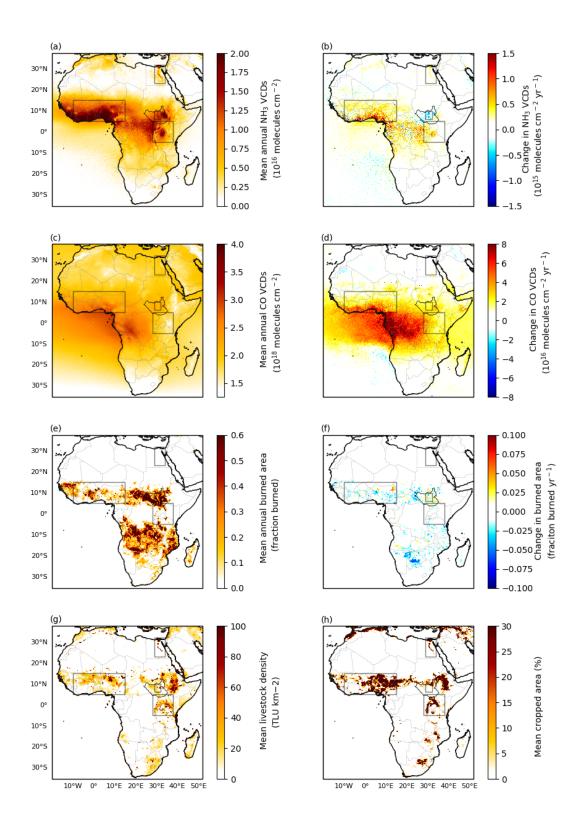
Linear trend analyses were conducted using linregress from the scipy.stats package
in Python v3.6.3. Statistical analyses of national scale data were conducted using ANOVA in
R. Data were log or rank transformed when necessary to meet the assumptions of ANOVA.
Values of α for treatment comparisons following significant ANOVA results were corrected
for multiple testing using Benjamini-Hochberg corrections.

- 236
- 237
- 238 **3. Results & Discussion**

3.1 Continental distributions and trends

240 Mean annual NH₃ concentrations for 2008-2018 are highest across the savannas 241 and forest-savanna mosaics in North equatorial Africa, and especially in West Africa; there 242 are smaller regional hotspots in the Lake Victoria basin, South Sudanese wetlands, and 243 along the Nile delta and river (Fig. 1a). Parts of these regions experience substantial 244 biomass burning (Fig. 1e), high livestock densities (Fig. 1g), and or high cropland cover 245 (Fig. 1h), all of which can contribute to NH₃ emissions. The high concentrations in West 246 Africa, which is one of the major global NH_3 hotspots (Van Damme et al., 2018), is likely the 247 result of biomass burning emissions. Biomass burning emissions tend to drive seasonal 248 variation in NH₃ VCDs in West Africa, with the largest emissions occurring late in the dry 249 season and early rainy season (Hickman et al., in review). In addition to local emissions, 250 biomass burning emissions and their reactive products are transported to the coast of West 251 Africa during both the northern hemisphere rainy season, when it is transported from 252 central and southern Africa, and during the dry season, when it is transported from 253 biomass burning regions to the east (Sauvage et al., 2007). Most areas with trends are 254 significant at P=0.2 or higher (Fig. S1).

255 In addition to being hotspots of mean NH₃ concentrations, some of these regions 256 have also experienced increases in NH_3 concentrations from 2008 to 2018 (Fig. 1b). Like 257 Warner et al. (2017) and Van Damme et al. (2021), we observed some increases in the 258 northern grasslands, central African forests, and the Nile region, but we also observe trends 259 in the Lake Victoria Basin, which Warner et al. (2017) did not, but Van Damme et al. (2021) 260 did. Also in contrast to Warner et al. (2017) but in line with Van Damme et al. (2021), we 261 observe a prominent decline in NH₃ VCDs over South Sudan (Fig. 1b, S1). We evaluate 262 several of these regions in more detail below.



264 Figure 1. Annual averages and trends in atmospheric NH₃ VCDs, CO VCDs, and burned 265 area, as well as spatial distribution of livestock density and cropped area across seven sub-266 Saharan African ecoregions. Mean annual (a) and trend (b) in atmospheric NH_3 VCDs from 267 IASI for the period 2008 through 2018. Mean annual (c) and trend (d) in annual 268 atmospheric CO VCDs from IASI for the same period. Mean annual (e) and trend (f) in 269 annual burned area from MODIS for 2008-2018. Livestock densities for 2007 from the FAO 270 (f), and mean cropped area from MODIS for 2008-2018 (g). The border of South Sudan is 271 highlighted in black, and several regions boxed: the Nile region at 30°N, the Sudd wetland 272 in South Sudan, the Lake Victoria region at the equator, and West Africa centered around 273 10°N.

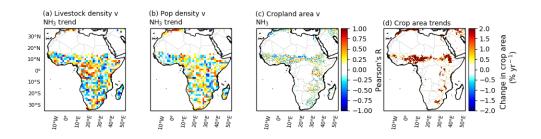


Figure 2. Relationships between NH₃ trends and livestock density, population density, and
cropland area, as well as changes in cropland area. Spatial correlations between changes in
annual atmospheric NH₃ VCDs and livestock density (a) and population density (b).
Correlation between cropland area and NH₃ VCDs for 2008 through 2018 (c). Change in
crop area for 2008 through 2018 (d). The NH₃ and crop area trends are based on data for
2008 through 2018, livestock density data are for the year 2007, population density data
are for the year 2017.

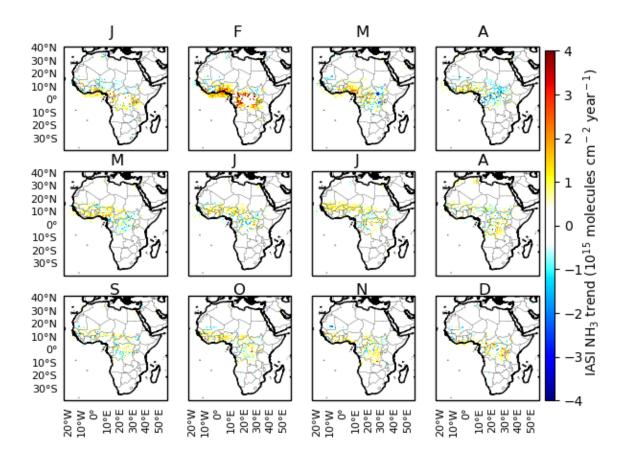


Figure 3. Change in mean monthly atmospheric NH₃ VCDs for the period 2008 through
 2018. Grid cells where mean annual NH₃ VCDs for the entire period are under 5x10¹⁵
 molecules cm⁻² are not displayed. Results significant at P=0.05 are presented in Figure S2.

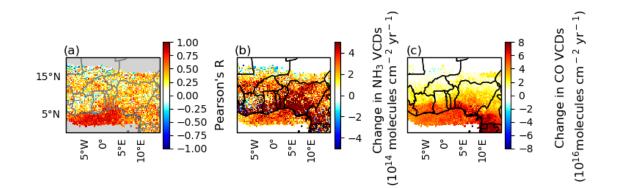


Figure 4. Correlation coefficient for the relationship between mean annual CO and NH₃ VCDs (a), changes in NH₃ VCDs (b) and changes in CO VCDs (c) over 2008 through 2018 in West Africa. Grid cells where mean annual NH₃ VCDs for the entire period are under 5x10¹⁵ molecules cm⁻² are not displayed. Results significant at P=0.05 for the entire continent are presented in Figure S5.

295

3.2 West Africa

297 The increasing trend in NH₃ VCDs over West Africa are centered over Nigeria and 298 the southern coast, and to a lesser extent across parts of the wet savanna (Fig. 1b). 299 Increases in NH₃ VCDs tend to be higher in grid cells with higher population densities in 300 Nigeria (Fig. 2b), suggesting a possible anthropogenic influence. The spatial distribution of 301 the mean annual NH₃ trend is overlapped by a substantial increase in mean annual CO 302 VCDs (Fig. 1b, 1d), pointing to a biomass burning source, as is also the case in central Africa. 303 Earlier studies have found substantial declines in annual burned area across the north 304 equatorial African biomass burning region as detected by MODIS (Andela et al., 2017; 305 Andela and Van Der Werf, 2014) and related declines in NO₂ VCDs across the region 306 (Hickman et al., in review; Hickman et al., 2021), which would seem to stand in contrast to the increasing CO and NH₃ trends observed here. 307

308 However, the annual decline in burned area and NO₂ VCDs is characterized by 309 heterogeneity when considering individual months. In West Africa, the dry season is 310 typically November to February or March. During the transition from the dry to rainy 311 season in February and March, NO₂ VCDs exhibit increasing rather than decreasing trends 312 in West Africa, though burned area patterns are not as clear when 2018 is included 313 (Hickman et al., 2021; Fig. S2, S3). Although these increases in NO_2 VCDs are small in the 314 annual context, they occur at a time of year when biomass burning combustion is less 315 complete, potentially due to greater fuel moisture and declining fire radiative power (Hickman et al., 2021; Zheng et al., 2018). These conditions would lead to greater 316

emissions of less oxidized species such as CO and NH₃, rather than the more fully oxidized
species such as CO₂ and NO₂ that dominate emissions during the peak of the biomass
burning season (Fig. S2, S4). Indeed, our observations suggest that much of the increasing
NH₃ trend occurs during this transitional period, with NH₃ VCDs increasing by roughly 7%
yr⁻¹ for all of Nigeria during February and March (Fig. 3; p<0.01). Variation in NH₃ VCDs
are positively correlated with CO VCDs (Fig. 4a, S5), which are also increasing during this
period (Fig. 4c, S4).

324 These correlations imply a biomass burning source for the increasing NH₃ VCDs in 325 West Africa; although the burned area trends are not as clear, it is important to remember 326 that MODIS undercounts burned area during this time of year by a factor of 3 to 6, and so 327 would be less sensitive to trends (Ramo et al., 2021; Roteta et al., 2019). Although there is 328 considerable gas flaring in Nigeria, gas flaring emissions have exhibited long-term negative 329 trends (Doumbia et al., 2019). In addition, although NO₂ VCDs were found to decrease 330 across the productive savannas of West Africa, regions of increasing NO_2 VCDs were 331 observed over large parts of Nigeria, further suggesting that there may be increases—or at 332 least smaller decreases—in biomass burning in the country (Hickman et al., 2021). It is 333 unlikely that changes in chemical sinks—specifically, the formation of nitrate aerosols in 334 reactions with NO_x or sulfate—are responsible for the increasing trend: the observed 335 increase in NO₂ VCDs observed during February and March would be expected to lead to a 336 shorter NH_3 lifetime and decreasing VCDs. In addition, emissions of SO_2 are relatively low 337 in West Africa, with moderate emissions occurring in Nigeria, but neither emissions nor 338 lifetime exhibit clear seasonal variation (Lee et al., 2011).

339 Small agricultural fires are likely an important contributor to the increasing NH₃ 340 VCDs during the dry-to-rainy season transitional period—a period when agricultural fires 341 are common in the region (Korontzi et al., 2006). There are large numbers of small fires 342 that are not detected by MODIS during these months: as noted above, estimates of burned 343 area during February, March, and April are revised upwards by roughly a factor of 3 to 6 344 over MODIS when small fires are included (Ramo et al., 2021; Roteta et al., 2019). Many of 345 these small fires are likely related to agricultural field preparation prior to planting 346 (Gbadegesin and Olusesi, 1994), which typically takes place in March or April (Vrieling et 347 al., 2011; Yegbemey et al., 2014). An increase in fires during this transitional period is also

consistent with one of the primary mechanisms behind the overall decline in burned area:
roughly half of the decline is attributed to increased population density and the expansion
of agricultural area, which contributes to the anthropogenic suppression of larger fires
(Andela et al., 2017; Andela and Van Der Werf, 2014). This agricultural expansion,
however, can be expected to be accompanied by increases in small fires used for the
removal of stubble or harvest byproduct (Gbadegesin and Olusesi, 1994), leading to the
increased emissions during the rainy-to-dry season transition observed here.

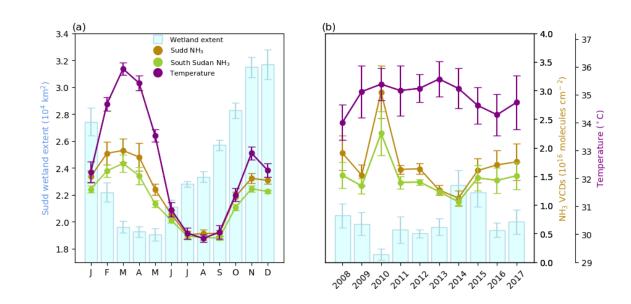
355 Globally, agricultural emissions from fertilized soils and livestock excreta are the 356 largest source of NH₃ (Bauer et al., 2016), and Warner et al. (Warner et al., 2017) suggest 357 that national-scale changes in fertilizer use could explain the NH₃ trend over Nigeria. 358 However, as noted above, much of the increase in West Africa occurs prior to the start of 359 the planting season—before fertilizer is applied—and appears likely to be due to biomass 360 burning emissions instead. Fertilizer or manure may make a contribution to the increasing 361 trend later in the year, as NH₃ VCDs increase in the wet savanna during May, June, and July 362 (Fig. 3), though there are also significant correlations between NH_3 and CO VCDs (Fig. 4), 363 suggesting that biomass burning may continue to play an important role. However, average 364 N fertilizer use in West Africa is universally under 40 kg N ha⁻¹ yr⁻¹, typically under 20 kg N 365 ha⁻¹ yr⁻¹, and is under 10 kg N ha⁻¹ yr⁻¹ in Nigeria—over an order of magnitude lower than 366 rates in Europe, the United States, and China (FAO, 2020). Although percentage changes in 367 fertilizer use are substantial, in absolute terms they represent increases of less than 2 kg N 368 ha⁻¹ yr⁻¹, and frequently less than 1 kg N ha⁻¹ yr⁻¹, a relatively small but perhaps not entirely 369 trivial perturbation to the N cycle. Between 2000 and 2007, total N deposition averaged 8.38 kg N ha⁻¹ yr⁻¹ in wet savanna and 14.75 kg N ha⁻¹ yr⁻¹ in forest ecosystems based on 370 371 surface sampling sites (Galy-Lacaux and Delon, 2014), and biological N fixation in tropical 372 and wet savannas has been estimated as ranging from 16 to 44 kg N ha⁻¹ yr⁻¹ (Bustamante 373 et al., 2006), suggesting that fertilizer increases may represent a 1 to 2% annual increase in 374 N inputs, But given the small magnitude of fertilizer applications, it appears unlikely that 375 changes in fertilizer use can explain the entirety of NH₃ increases during the growing 376 season.

3.3. South Sudan

379 The most notable declining trend in NH₃ VCDs occurs in South Sudan over the Sudd 380 wetlands at a rate of over 2% yr⁻¹ (Fig. 1b; p=0.20). It appears that this decline is related to 381 interannual variation in the flooded extent of the Sudd, a vast wetland that connects the 382 White and Blue Nile tributaries. Seasonal variation of inflow to the Sudd leads to variation 383 in flooded extent: an area of roughly 15,000 km² is permanently flooded, and another 384 roughly 15,000 km² is temporarily flooded each year, with considerable interannual 385 variation in the total flooded area (Di Vittorio and Georgakakos, 2018). Among other 386 factors, drying soils should increase production and emissions of NH₃ from soils, as Eq. 1 is 387 shifted to the right (Clarisse et al., 2019). Earlier work evaluating an NH₃ hotspot over Lake 388 Natron in Tanzania found that the drying of seasonally flooded soils leads to large 389 emissions of NH₃: As the waters of Lake Natron recede during the dry season each year 390 and the surrounding mud flats dry out, NH₃ VCDs increase rapidly, with hotspots appearing 391 over the mudflats (Clarisse et al., 2019). These elevated VCDs are attributed to multiple 392 possible factors, including the effects of drying on concentrations of NH₃ in solution (which 393 increases the concentration gradient with the atmosphere), reduced biological uptake of 394 NH₃, convective transport of dissolved NH₃ from depth to the soil surface, and increased 395 mineralization of labile organic matter (Clarisse et al., 2019).

396 We find the same clear seasonal relationship between wetland flooded extent and 397 NH₃ concentrations over the Sudd—VCDs increase as waters recede from the temporarily 398 flooded area, leading to annual maxima from February through May (Fig. 5a; bounding box 399 of 29E to 31.5E and 6N to 9.9N). Like the entire country, seasonal variation in NH_3 VCDs 400 over the Sudd follow variation in surface temperature, but NH₃ concentrations over the 401 Sudd are substantially elevated compared to surrounding regions during this time of year 402 but not others, suggesting that an mechanism in addition to temperature is contributing to 403 the elevated emissions in the Sudd during February through May, a period that spans the 404 end of the dry season and start of the rainy season (Fig. S6). This conclusion is supported 405 by an analysis of interannual variation of VCDs during the February through May period: 406 Interannual variation in NH₃ VCDs is largely decoupled from variation in temperature, but

- NH₃ VCDs appear to vary inversely with the amount of area that dries out each year (Fig. 5b). Over the period for which flooded extent data are currently available for the Sudd, the minimum flooded extent tends to increase—that is, less area dries out each year—resulting in an overall decline in NH₃ VCDs. Linear regression reveals that this change in flooded extent explains a large proportion of the annual variation in NH_3 in the Sudd bounding box (r=-0.65, p=0.04), as well as for the country as a whole (r=-0.63, p=0.05). These analyses strongly suggest that the declining trend in NH₃ over the Sudd is a direct result of an overall increase in the minimum flooded extent over the observation period.





418 Figure 5. Mean (a) monthly and (b) February through May annual mean flooded
419 extent of the Sudd, surface temperatures over South Sudan, and NH₃ VCDs over the Sudd
420 and the entirety of South Sudan for the period 2008 through 2017.

It is possible that conflict in South Sudan could contribute to the decline in NH₃
VCDs. In 2013, a civil conflict emerged in South Sudan that was ultimately responsible for
the displacement of millions of people (Global Internal Displacement Monitoring Centre,
2020; World Bank, 2019) and the disruption of livestock migration patterns (Idris, 2018).
However, these disruptions appeared only after the onset of the long-term change in NH₃,
and appear unlikely to make an important contribution to the observed interannual
variation (SI Text, Fig. S7, S8).

431 It is unlikely that changes in chemical sinks are responsible for the decline in NH₃ 432 VCDs. VCDs of tropospheric NO₂ are also decreasing in the region (Fig. S9), which is 433 suggestive of less formation of particulate-phase ammonium rather than more. 434 Anthropogenic SO₂ emissions in Africa in general and South Sudan in particular are very 435 low (European Commission Joint Research Centre (JRC)/Netherlands Environmental 436 Assessment Agency (PBL), 2016), and would not be expected to be emitted from the Sudd; 437 more generally, the clear spatial association between the NH_3 trend and the Sudd (Fig. 1, 438 Fig. S10) is strongly suggestive of changes in emissions rather than atmospheric processes 439 being responsible for the trend.

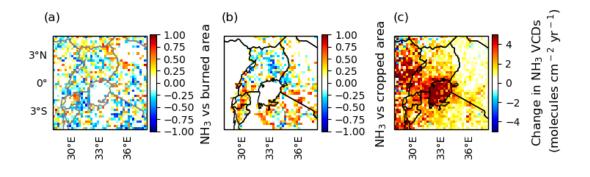
440

441

3.4 Lake Victoria Basin region

442 The Lake Victoria Basin and its surroundings—an area including elevated mean NH₃ 443 VCDs—exhibit an increasing NH₃ trend (Fig. 1b, Fig. 7, Fig. S11), which appears to be the 444 result of increasing agricultural activity in the area. The region includes a high and 445 increasing density of agricultural land (Fig. 1h, Fig. 2d, Fig. S12), and these increases in 446 cropped area are positively correlated with increases in NH₃ VCDs across much of the 447 region (Fig. 2c). The northern and southern halves of the Lake Victoria region have distinct 448 growing seasons: in the north, the season generally starts in April, whereas in the south, it 449 starts in November or December (Vrieling et al., 2011). The long-term trend reflects this 450 seasonality, with increases in the north and south occurring during their respective 451 growing seasons (Fig. 3). Fertilizer use in the Lake Victoria region is low: national averages 452 range from about 1 to 3 kg nutrients ha⁻¹ yr⁻¹ in Uganda to about 35 to 40 kg nutrients ha⁻¹ 453 yr⁻¹ in Kenya (Elrys et al., 2019; World Bank, 2019); to put these numbers in context, 454 Organization for Economic Cooperation and Development (OECD) countries use about 135-455 140 kg nutrients ha⁻¹ yr⁻¹ (World Bank, 2019). Although rates of fertilizer use have 456 increased by substantial proportions, the absolute amount of increase is relatively small, 457 typically roughly 1 to 10 kg nutrients decade⁻¹. Unlike in West Africa, however, interannual 458 variation in burned area (Fig. S13) does not exhibit a clear relationship with changes in 459 NH_3 VCDs. Consequently, we expect that both the expansion and intensification of

460 agriculture in the region contribute to the increasing NH₃ VCDs.



461

462 Figure 7. Changes in NH₃ VCDs and their relationship with burned area and
463 cropped area over the Lake Victoria region for the 2008 through 2018 period. (a)
464 Correlation coefficients for the relationship between NH₃ VCDs and burned area. (b)

465 Correlation coefficients for the relationship between NH₃ VCDs and cropped area, including

 $466 \qquad mosaics \ of \ crops \ and \ natural \ vegetation \ cover. \ (c) \ Changes \ in \ NH_3 \ VCDs.$

467

468

3.5 National-scale relationships

Examining relationships at a national scale can provide insight into relationships
between changes in agricultural or biomass burning and changes in atmospheric NH₃ VCDs
at larger scales. When grouping countries into three bins based on their annual percentage
changes in NH₃ VCDs, there is some evidence for a broad relationship between livestock
and NH₃ VCDs at the national scale (Fig. 8). The rate of change in national-scale NH₃ VCDs
varies significantly among bins (p<0.001; rank transformed). The annual percentage

- 475 changes in livestock in TLUs vary significantly by bin (p=0.056), with the middle bin higher
- 476 than the bottom bin (p=0.03) and the high bin higher than the bottom bin, though not
- 477 significantly so (p=0.21). , Annual percentage changes in fertilizer N (p=0.23) and crop
- 478 production (p=0.62; rank transformed) did not vary by bin.
- 479

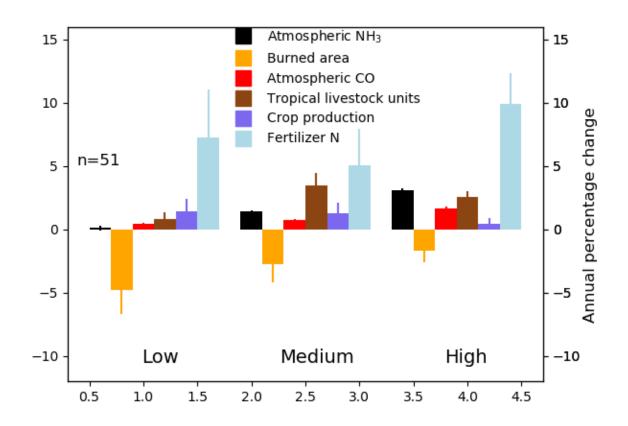


Figure 8. Annual percentage changes in national mean annual NH₃ VCDs, burned area, CO
VCDs, livestock, crop yield, and fertilizer N use for African countries with low, medium, or
high rates of NH₃ VCD change. Error bars represent the standard error of the mean. See
Table S1 for the list of countries in each bin and Fig. S14 for an expanded set of variables.

486

487 Instead of a direct agricultural relationship with changes in NH₃ VCDs, there is the

488 possibility that changes in biomass burning are associated with changes in NH₃ VCDs.

489 Although the differences in the annual percentage change in burned area were not

490 significant among bins (p=0.48; rank transformed), the overall pattern is consistent with 491 earlier results finding that a reduction in burned area across the northern biomass burning 492 region was associated in part with the expansion of agriculture and presumed 493 anthropogenic suppression of fire (Andela et al., 2017; Andela and Van Der Werf, 2014). 494 However, burned area as measured by MODIS is likely an imperfect predictor for NH₃ 495 emissions—as noted previously, MODIS underestimates burned area by a factor of 3 to 6 496 during shoulder seasons (Roteta et al., 2019), which is when fires are expected to emit 497 more reduced species such as NH_3 (Zheng et al., 2018). In contrast to burned area, the 498 annual change in column densities of CO—which tends to be co-emitted with NH₃ from 499 fires—differed significantly among bins (p<0.001; rank transformed) and was significantly 500 higher in the high bin than in the low or medium bins. (p<0.001, post-hoc tests). The 501 higher annual CO changes in the high bin could related to larger anthropogenic fossil fuel 502 emissions, but we see no difference among bins in growth rates of CO_2 emissions (p=0.45; 503 Figure S14); such a difference would be expected if differences in economic development 504 were responsible for the CO differences. These results leave open the possibility that 505 changes in either biofuel emissions or biomass burning emissions—perhaps from smaller 506 fires not observed in the MODIS burned area product—may be primarily responsible for 507 the difference in CO between bins, and may be contributing to the differences in NH₃ 508 between bins. Changes in NO₂ VCDs and SO₂ concentrations can affect the lifetime of NH₃ 509 (the latter by changing SO_4 concentrations), but do not appear to make an important 510 contribution to the observed trends in NH₃ VCDs among bins (Fig. S14, SI text). 511 Temperature, likewise, does not appear to play an important role (SI text).

- 512
- 513

4. Conclusion

514 Using IASI, we have observed both increases and decreases in atmospheric NH₃ 515 VCDs in different regions in Africa between 2008 and 2018, with different factors affecting 516 trends in different regions.

517 We observed increases in NH₃ VCDs in West Africa, which earlier work had 518 concluded was likely related to increased fertilizer use. Fertilizer is not typically applied in 519 West Africa until the start of the growing season—often April—but we find that much of 520 the NH₃ increase occurs during February and March, suggesting that increasing fertilizer use is unlikely to provide a complete explanation for the NH₃ trend. Agriculture may
nevertheless play a role, with enhanced burned area and especially CO concentrations in
February suggestive of increased burning of crop stubble in preparation for planting during
this time of year. Fires in this region tend to emit a greater proportion of less oxidized
species such as NH₃ at the end of the dry season, consistent with a biomass burning source
for the increasing NH₃ VCDs.

527 Decreases in NH₃ VCDs were largest in South Sudan, especially over the Sudd 528 wetland, where NH_3 VCDs vary seasonally with the extent of area flooded. As the 529 temporarily flooded areas of the Sudd dry out each year, NH₃ VCDs increase as reduction in 530 soil moisture drives increased production and volatilization of NH₃. The area of the Sudd 531 that is flooded each year varies, and from 2008 until 2015, the area that remains flooded 532 during the dry season generally increased, producing a positive overall trend for the period 533 of 2008 through 2017. This increase in the dry season flooded area drove a decrease in 534 NH₃ VCDs: with less soil drying out, the seasonal maxima in NH₃ VCDs were lower. 535 Although it is possible that conflict in South Sudan could contribute to changes in NH₃ 536 VCDs, the timing and distribution of conflict events and human displacement suggest that 537 other factors are likely more important.

538 Modest increases in NH₃ VCDs were observed in the Lake Victoria region. This 539 region has experienced increases in agricultural area during the IASI observation period, 540 and these changes explained a large proportion of the variation in NH₃ VCDs across large 541 patches of the region, where biomass burning could not. We expect that both expansion 542 and intensification of agriculture in this region could contribute to the positive NH₃ trend.

543 Considering national-scale statistics, comparisons between equally sized bins of 17 544 countries each suggested that changes in biomass burning emissions and livestock 545 emissions could contribute to differences in NH₃ VCDs among countries, but variables 546 related to cropped agriculture such as cropped area or fertilizer N use did not appear to be 547 important factors at this scale. This may be because although fertilizer use has been 548 increasing in sub-Saharan Africa, it remains extremely low relative to other continents, and 549 relative to the levels needed to attain food security. Average fertilizer use in most 550 countries in the region is under 20 kg N ha⁻¹ yr⁻¹, and sometimes less than 5 kg N ha⁻¹ yr⁻¹. 551 Although recommended fertilizer rates are lower in most African countries than in the U.S.

552 or Europe, increasing N inputs to 50, 100, or 150 kg N ha⁻¹ yr⁻¹ would represent a major 553 perturbation to the regional N cycle, and potentially a large new source of NH₃ to the 554 atmosphere. West Africa is already a global NH₃ hotspot (Van Damme et al., 2018), 555 suggesting that encouraging policies that can help to limit NH₃ emissions during the early 556 stages of agricultural intensification in Africa may help mitigate potential impacts on the 557 atmosphere. Fortunately, agricultural practices such as sub-surface application of 558 fertilizer, which is already being promoted to smallholder farmers, can serve to both limit 559 NH₃ emissions also help to increase crop yields.

560 These past and anticipated future trends also make the case for expanding capacity 561 for atmospheric monitoring in sub-Saharan Africa. Although long-term monitoring 562 networks have been established in West Africa (Adon et al., 2010; Ossohou et al., 2019) and 563 South Africa (Conradie et al., 2016) as part of the INDAAF network, it is mainly focused on 564 deposition, the spatio-temporal resolution of surface measurements is very coarse when 565 compared to the data available in other parts of the world, and will limit our ability to 566 understand how agricultural and socio-economic development in Africa affect the 567 atmosphere. Satellite observations can help to bridge some of these data gaps, but have 568 their own spatio-temporal limitations, and would further benefit from additional high-569 quality surface observations for evaluation of retrieval products.

570

571 Data availability: All data used in this study are available from public sources, with the 572 exception of Sudd wetland extent, which is available by request from Courtney Di Vittorio. The 573 IASI NH₃ and CO data are available from The IASI https://iasi.aeris-data.fr. The NOAA Global 574 Surface is available https://data.nodc.noaa.gov/cgi-Temperature Dataset at 575 bin/iso?id=gov.noaa.ncdc:C01585. MODIS burned area data are available from 576 https://www.globalfiredata.org/data.html. MODIS agricultural area are available at 577 https://lpdaac.usgs.gov/products/mcd12c1v006/. TRMM 3B42 precipitation data are available 578 from https://pmm.nasa.gov/data-access/downloads/trmm. The Gridded Livestock of the World 579 data are available from https://livestock.geo-wiki.org/home-2/. Population density data for 2017 580 are available at https://landscan.ornl.gov/downloads/2017. FAO national crop production and 581 fertilizer N data are available at http://www.fao.org/faostat/en/. Data on conflict events from

- 582 ACLED are available at https://acleddata.com/#/dashboard. World Bank national statistics
- 583 on refugees and internally displaced people are available at https://data.worldbank.org.

- Author Contribution: J.E.H. designed the study, conducted the analysis, and wrote the paper.
 NA, ED, CD, MO, CG-L, KT, and SEB contributed to study design and edited the paper. LC, P-
- 587 FC, and MVD developed the original IASI trace gas retrievals and edited the paper.
- 588 The authors declare that they have no conflict of interest.
- 589

590 **References**

- 591 van der A, R. J., Eskes, H. J., Boersma, K. F., van Noije, T. P. C., Van Roozendael, M., De Smedt,
- 592 I., Peters, D. H. M. U. and Meijer, E. W.: Trends, seasonal variability and dominant NOx
- source derived from a ten year record of NO2 measured from space, J. Geophys. Res.
- 594 Atmos., 113(4), D04302, doi:10.1029/2007JD009021, 2008.
- Adon, M., Galy-Lacaux, C., Yoboué, V., Delon, C., Lacaux, J. P., Castera, P., Gardrat, E., Pienaar,
- 596 J., Al Ourabi, H., Laouali, D., Diop, B., Sigha-Nkamdjou, L., Akpo, a., Tathy, J. P., Lavenu, F. and
- 597 Mougin, E.: Long term measurements of sulfur dioxide, nitrogen dioxide, ammonia, nitric
- acid and ozone in Africa using passive samplers, Atmos. Chem. Phys., 10(15), 7467–7487,
- 599 doi:10.5194/acp-10-7467-2010, 2010.
- 600 AGRA: AGRA in 2008: Building on the New Momentum in African Agriculture., 2009.
- 601 Andela, N. and Van Der Werf, G. R.: Recent trends in African fires driven by cropland
- expansion and El Niño to la Niña transition, Nat. Clim. Chang., 4(9), 791–795,
- 603 doi:10.1038/nclimate2313, 2014.
- Andela, N., Morton, D. C., Giglio, L., Chen, Y., Van Der Werf, G. R., Kasibhatla, P. S., DeFries, R.
- 605 S., Collatz, G. J., Hantson, S., Kloster, S., Bachelet, D., Forrest, M., Lasslop, G., Li, F., Mangeon,
- 606 S., Melton, J. R., Yue, C. and Randerson, J. T.: A human-driven decline in global burned area,
- 607 Science (80-.)., 356(6345), 1356–1362, doi:10.1126/science.aal4108, 2017.
- Anker, H. T., Baaner, L., Backes, C., Keessen, A. and Möckel, S.: Comparison of ammonia
- 609 regulation in Germany , the Netherlands and Denmark legal framework, Copenhagen.,
- 610 2018.

- 611 Bauer, S. E., Tsigaridis, K. and Miller, R.: Significant atmospheric aerosol pollution caused by
- 612 world food cultivation, Geophys. Res. Lett., 43(10), 5394–5400,
- 613 doi:10.1002/2016GL068354, 2016.
- 614 Behera, S. N., Sharma, M., Aneja, V. P. and Balasubramanian, R.: Ammonia in the
- 615 atmosphere: A review on emission sources, atmospheric chemistry and deposition on
- 616 terrestrial bodies, Environ. Sci. Pollut. Res., 20(11), 8092–8131, doi:10.1007/s11356-013-
- 617 2051-9, 2013.
- Bouwman, A. F., Lee, D. S., Asman, W. A. H., Dentener, F. J., Van Der Hoek, K. W. and Olivier, J.
- 619 G. J.: A global high-resolution emission inventory for ammonia, Global Biogeochem. Cycles,
- 620 11(4), 561–587, 1997.
- 621 Bustamante, M. M. C., Medina, E., Asner, G. P., Nardoto, G. B. and Garcia-Montiel, D. C.:
- Nitrogen cycling in tropical and temperate savannas, Biogeochemistry, 79(1–2), 209–237,
- 623 doi:10.1007/s10533-006-9006-x, 2006.
- 624 Cahoon, D. R., Stocks, B. J., Levine, J. S., Cofer, W. R. and O'Neill, K. P.: Seasonal distribution of
- 625 African savanna fires, Nature, 359(6398), 812–815, doi:10.1038/359812a0, 1992.
- 626 Chen, L.-W. A., Verburg, P., Shackelford, A., Zhu, D., Susfalk, R., Chow, J. C. and Watson, J. G.:
- 627 Moisture effects on carbon and nitrogen emission from burning of wildland biomass,
- 628 Atmos. Chem. Phys., 10(14), 6617–6625, doi:10.5194/acp-10-6617-2010, 2010.
- 629 Chilonda, P. and Otte, J.: Indicators to monitor trends in livestock production at national,
- 630 regional and international levels, Livest. Res. Rural Dev., 18(8), 1–12, 2006.
- 631 Clarisse, L., Clerbaux, C., Dentener, F., Hurtmans, D. and Coheur, P. F.: Global ammonia
- distribution derived from infrared satellite observations, Nat. Geosci., 2(7), 479–483,
- 633 doi:10.1038/ngeo551, 2009.
- 634 Clarisse, L., Van Damme, M., Gardner, W., Coheur, P.-F., Clerbaux, C., Whitburn, S., Hadji-
- 635 Lazaro, J. and Hurtmans, D.: Atmospheric ammonia (NH3) emanations from Lake Natron's
- 636 saline mudflats, Sci. Rep., 9, 4441, doi:10.1038/s41598-019-39935-3, 2019.
- 637 Conradie, E. H., Van Zyl, P. G., Pienaar, J. J., Beukes, J. P., Galy-Lacaux, C., Venter, A. D. and
- 638 Mkhatshwa, G. V.: The chemical composition and fluxes of atmospheric wet deposition at
- 639 four sites in South Africa, Atmos. Environ., 146, 113–131,
- 640 doi:10.1016/j.atmosenv.2016.07.033, 2016.
- 641 Van Damme, M., Wichink Kruit, R. J., Schaap, M., Clarisse, L., Clerbaux, C., Coheur, P. F.,

- 642 Dammers, E., Dolman, A. J. and Erisman, J. W.: Evaluating 4 years of atmospheric ammonia
- 643 (NH3) over Europe using IASI satellite observations and LOTOS-EUROS model results, J.

644 Geophys. Res., 119(15), 9549–9566, doi:10.1002/2014JD021911, 2014a.

- 645 Van Damme, M., Clarisse, L., Heald, C. L., Hurtmans, D., Ngadi, Y., Clerbaux, C., Dolman, A. J.,
- 646 Erisman, J. W. and Coheur, P. F.: Global distributions, time series and error characterization
- 647 of atmospheric ammonia NH3 from IASI satellite observations, Atmos. Chem. Phys., 14(6),
- 648 2905–2922, doi:10.5194/acp-14-2905-2014, 2014b.
- 649 Van Damme, M., Clarisse, L., Dammers, E., Liu, X., Nowak, J. B., Clerbaux, C., Flechard, C. R.,
- 650 Galy-Lacaux, C., Xu, W., Neuman, J. A., Tang, Y. S., Sutton, M. A., Erisman, J. W. and Coheur, P.
- F.: Towards validation of ammonia (NH3) measurements from the IASI satellite, Atmos.
- 652 Meas. Tech., 8(3), 1575–1591, doi:10.5194/amt-8-1575-2015, 2015.
- 653 Van Damme, M., Whitburn, S., Clarisse, L., Clerbaux, C., Hurtmans, D. and Coheur, P. F.:
- 654 Version 2 of the IASI NH3 neural network retrieval algorithm: Near-real-time and
- reanalysed datasets, Atmos. Meas. Tech., 10(12), 4905–4914, doi:10.5194/amt-10-49052017, 2017.
- 657 Van Damme, M., Clarisse, L., Whitburn, S., Hadji-Lazaro, J., Hurtmans, D., Clerbaux, C. and
- 658 Coheur, P.: Industrial and agricultural ammonia point sources exposed, Nature, 564, 99–
- 659 103, 2018.
- 660 Van Damme, M., Clarisse, L., Franco, B., Sutton, M. A., Erisman, J. W., Wichink Kruit, R., van
- 661 Zanten, M., Whitburn, S., Hadji-Lazaro, J., Hurtmans, D., Clerbaux, C. and Coheur, P.-F.:
- 662 Global, regional and national trends of atmospheric ammonia derived from a decadal
- 663 (2008-2018) satellite record, Environ. Res. Lett., 2021.
- 664 Dammers, E., Palm, M., Van Damme, M., Vigouroux, C., Smale, D., Conway, S., Toon, G. C.,
- Jones, N., Nussbaumer, E., Warneke, T., Petri, C., Clarisse, L., Clerbaux, C., Hermans, C.,
- 666 Lutsch, E., Strong, K., Hannigan, J. W., Nakajima, H., Morino, I., Herrera, B., Stremme, W.,
- 667 Grutter, M., Schaap, M., Kruit, R. J. W., Notholt, J., Coheur, P. F. and Erisman, J. W.: An
- 668 evaluation of IASI-NH3 with ground-based Fourier transform infrared spectroscopy
- 669 measurements, Atmos. Chem. Phys., 16(16), 10351–10368, doi:10.5194/acp-16-10351-
- 670 **2016**, 2016.
- 671 Denier Van Der Gon, H. and Bleeker, A.: Indirect N2O emission due to atmospheric N
- deposition for the Netherlands, Atmos. Environ., 39(32), 5827–5838,

- 673 doi:10.1016/j.atmosenv.2005.06.019, 2005.
- Dobson, J. E., Bright, E. A., Coleman, P. R., Durfee, R. C. and Worley, B. A.: A global population
- database for estimating populations at risk, Photogramm. Eng. Remote Sens., 66(7), 849–
- 676 857 [online] Available from: In Book..., 2000.
- 677 Doumbia, E. H. T., Liousse, C., Keita, S., Granier, L., Granier, C., Elvidge, C. D., Elguindi, N. and
- 678 Law, K.: Flaring emissions in Africa: Distribution, evolution and comparison with current
- 679 inventories, Atmos. Environ., 199(November 2018), 423–434,
- 680 doi:10.1016/j.atmosenv.2018.11.006, 2019.
- 681 Elrys, A. S., Abdel-Fattah, M. K., Raza, S., Chen, Z. and Zhou, J.: Spatial trends in the nitrogen
- budget of the African agro-food system over the past five decades, Environ. Res. Lett.,
- 683 14(12), doi:10.1088/1748-9326/ab5d9e, 2019.
- 684 European Commission Joint Research Centre (JRC)/Netherlands Environmental
- 685 Assessment Agency (PBL): Emission Database for Global Atmospheric Research (EDGAR),
- 686 release version 4.3.1, 2016.
- 687 Fan, Y. and van den Dool, H.: A global monthly land surface air temperature analysis for
- 688 1948-present, J. Geophys. Res. Atmos., 113(1), D01103, doi:10.1029/2007JD008470, 2008.
- 689 FAO: FAO Statistics Database, FAOSTAT Stat. Database [online] Available from:
- 690 http://www.fao.org/faostat/en/ (Accessed 1 January 2019), 2020.
- 691 Friedl, M. A., McIver, D. K., Hodges, J. C. F., Zhang, X. Y., Muchoney, D., Strahler, A. H.,
- 692 Woodcock, C. E., Gopal, S., Schneider, A., Cooper, A., Baccini, A., Gao, F. and Schaaf, C.: Global
- 693 land cover mapping from MODIS: Algorithms and early results, Remote Sens. Environ.,
- 694 83(1-2), 287-302, doi:10.1016/S0034-4257(02)00078-0, 2002.
- 695 Galy-Lacaux, C. and Delon, C.: Nitrogen emission and deposition budget in West and Central
- 696 Africa, Environ. Res. Lett., 9(12), doi:10.1088/1748-9326/9/12/125002, 2014.
- 697 Gbadegesin, A. and Olusesi, B. B.: Effects of land clearing methods on soil physical and
- 698 hydrological properties in southwestern Nigeria, Environmentalist, 14(4), 297–303
- 699 [online] Available from: http://www.scopus.com/inward/record.url?eid=2-s2.0-
- 700 0022858519&partnerID=40&md5=3051b1dd91ff2990e37fe7466872923e, 1994.
- 701 George, M., Clerbaux, C., Hurtmans, D., Turquety, S., Coheur, P. F., Pommier, M., Hadji-
- 702 Lazaro, J., Edwards, D. P., Worden, H., Luo, M., Rinsland, C. and McMillan, W.: Carbon
- 703 monoxide distributions from the IASI/METOP mission: Evaluation with other space-borne

- remote sensors, Atmos. Chem. Phys., 9(21), 8317–8330, doi:10.5194/acp-9-8317-2009,
 2009.
- Giglio, L., Boschetti, L., Roy, D. P., Humber, M. L. and Justice, C. O.: The Collection 6 MODIS
- burned area mapping algorithm and product, Remote Sens. Environ., 217, 72–85,
- 708 doi:10.1016/j.rse.2018.08.005, 2018.
- 709 Global Internal Displacement Monitoring Centre: Global Internal Displacement Database,
- 710 [online] Available from: https://www.internal-displacement.org/database/displacement-
- 711 data, 2020.
- 712 Goode, J. G., Yokelson, R. J., Susott, R. A. and Ward, D. E.: Trace gas emissions from
- 713 laboratory biomass fires measured by open-path Fourier transform infrared spectroscopy:,
- 714 J. Chem. Inf. Model., 104(D17), 21237–21245, 1999.
- 715 Guo, X., Clarisse, L., Wang, R., Van Damme, M., Whitburn, S., Coheur, P., Clerbaux, C., Franco,
- 716 B., Pan, D., Golston, L. M., Wendt, L., Sun, K., Tao, L., Miller, D., Mikoviny, T., Müller, M.,
- 717 Wisthaler, A., Tevlin, A. G., Murphy, J. G., Nowak, J. B., Roscioli, J. R., Volkamer, R., Kille, N.,
- 718 Neuman, J. A., Eilerman, S. J., Crawford, J. H., Yacovitch, T. I., Barrick, J. D., Scarino, A. J. and
- 719 Zondlo, M. A.: Validation of IASI satellite ammonia observations at the pixel scale using in-
- situ vertical profiles, J. Geophys. Res. Atmos., 126, e2020JD033475,
- 721 doi:10.1029/2020jd033475, 2021.
- Hazell, P. and Wood, S.: Drivers of change in global agriculture., Philos. Trans. R. Soc. Lond.
- 723 B. Biol. Sci., 363(1491), 495–515, doi:10.1098/rstb.2007.2166, 2008.
- Hickman, Jonathan, E., Andela, N., Galy-Lacaux, C., Ossohou, M., Dammers, E., Van Damme,
- 725 M., Clarisse, L., Tsigaridis, K. and Bauer, S. E.: Continental and ecoregion-specific drivers of
- atmospheric NO2 and NH3 seasonality over Africa revealed by satellite observations,
- 727 Global Biogeochem. Cycles, in review.
- Hickman, J. E., Dammers, E., Galy-Lacaux, C. and Van Der Werf, G. R.: Satellite evidence of
- substantial rain-induced soil emissions of ammonia across the Sahel, Atmos. Chem. Phys.,
- 730 18(22), 16713–16727, doi:10.5194/acp-18-16713-2018, 2018.
- 731 Hickman, J. E., Andela, N., Tsigaridis, K., Galy-Lacaux, C., Ossohou, M. and Bauer, S. E.:
- 732 Reductions in NO 2 burden over north equatorial Africa from decline in biomass burning in
- 733 spite of growing fossil fuel use, 2005 to 2017, Proc. Natl. Acad. Sci., 118(7), e2002579118,
- 734 doi:10.1073/pnas.2002579118, 2021.

- Huffman, G. J., Adler, R. F., Bolvin, D. T., Gu, G., Nelkin, E. J., Bowman, K. P., Hong, Y., Stocker,
- 736 E. F. and Wolff, D. B.: The TRMM Multisatellite Precipitation Analysis (TMPA): Quasi-Global,
- 737 Multivear, Combined-Sensor Precipitation Estimates at Fine Scales, J. Hydrometeorol., 8(1),
- 738 38–55, doi:10.1175/jhm560.1, 2007.
- 739 Hurtmans, D., Coheur, P. F., Wespes, C., Clarisse, L., Scharf, O., Clerbaux, C., Hadji-Lazaro, J.,
- 740 George, M. and Turquety, S.: FORLI radiative transfer and retrieval code for IASI, J. Quant.
- 741 Spectrosc. Radiat. Transf., 113(11), 1391–1408, doi:10.1016/j.jqsrt.2012.02.036, 2012.
- 742 Idris, I.: Livestock and conflict in South Sudan K4D Helpdesk Report 484, Brighton.
- 743 [online] Available from:
- 744 https://assets.publishing.service.gov.uk/media/5c6abdec40f0b61a22792fd5/484_Livest
- 745 ock_and_Conflict_in_South_Sudan.pdf, 2018.
- 746 Kanter, D. R.: Nitrogen pollution: a key building block for addressing climate change, Clim.
- 747 Change, 147(1–2), 11–21, doi:10.1007/s10584-017-2126-6, 2018.
- 748 Kerzenmacher, T., Dils, B., Kumps, N., Blumenstock, T., Clerbaux, C., Coheur, P. F., Demoulin,
- P., García, O., George, M., Griffith, D. W. T., Hase, F., Hadji-Lazaro, J., Hurtmans, D., Jones, N.,
- 750 Mahieu, E., Notholt, J., Paton-Walsh, C., Raffalski, U., Ridder, T., Schneider, M., Servais, C. and
- 751 De Mazière, M.: Validation of IASI FORLI carbon monoxide retrievals using FTIR data from
- 752 NDACC, Atmos. Meas. Tech., 5(11), 2751–2761, doi:10.5194/amt-5-2751-2012, 2012.
- 753 Korontzi, S., McCarty, J., Loboda, T., Kumar, S. and Justice, C.: Global distribution of
- agricultural fires in croplands from 3 years of Moderate Resolution Imaging
- 755 Spectroradiometer (MODIS) data, Global Biogeochem. Cycles, 20(2), GB2021,
- 756 doi:10.1029/2005GB002529, 2006.
- 757 Krupa, S. V.: Effects of atmospheric ammonia (NH3) on terrestrial vegetation: A review,
- 758 Environ. Pollut., 124(2), 179–221, doi:10.1016/S0269-7491(02)00434-7, 2003.
- Lee, C., Martin, R. V., Van Donkelaar, A., Lee, H., Dickerson, R. R., Hains, J. C., Krotkov, N.,
- 760 Richter, A., Vinnikov, K. and Schwab, J. J.: SO2 emissions and lifetimes: Estimates from
- 761 inverse modeling using in situ and global, space-based (SCIAMACHY and OMI)
- 762 observations, J. Geophys. Res. Atmos., 116(6), 1–13, doi:10.1029/2010JD014758, 2011.
- 763 Lelieveld, J., Evans, J. S., Fnais, M., Giannadaki, D. and Pozzer, A.: The contribution of outdoor
- air pollution sources to premature mortality on a global scale, Nature, 525(7569), 367–371,
- 765 doi:10.1038/nature15371, 2015.

- Nicholson, S., Some, B., McCollum, J., Nelkin, E., Klotter, D., Berte, Y., Diallo, B., Gaye, I.,
- 767 Kpabeba, G., Ndiaye, O., Noukpozounkou, J., Tanu, M., Thiam, A., Toure, A. and Traore, A.:
- 768 Validation of TRMM and other rainfall estimates with a high-density gauge dataset for West
- 769 Africa. Part II: Validation of TRMM rainfall products, J. Appl. Meteorol., 42, 1355–1368,
- 770 doi:Article, 2003.
- 771 Ossohou, M., Galy-Lacaux, C., Yoboué, V., Hickman, J. E., Gardrat, E., Adon, M., Darras, S.,
- T72 Laouali, D., Akpo, A., Ouafo, M., Diop, B. and Opepa, C.: Trends and seasonal variability of
- atmospheric NO2 and HNO3 concentrations across three major African biomes inferred
- from long-term series of ground-based and satellite measurements, Atmos. Environ., 207,
- 775 148–166, 2019.
- Pamela A. Matson, McDowell, W. H., Townsend, A. R. and Vitousek, P. M.: The globalization
- of N deposition: ecosystem consequences in tropical environments, Biogeochemistry, 46,
- 778 67-83 [online] Available from: http://academic.engr.arizona.edu/HWR/Brooks/GC572-
- 779 2004/readings/matson.pdf, 1999.
- 780 Pommier, M., Law, K. S., Clerbaux, C., Turquety, S., Hurtmans, D., Hadji-Lazaro, J., Coheur, P.
- F., Schlager, H., Ancellet, G., Paris, J. D., Nédlec, P., Diskin, G. S., Podolske, J. R., Holloway, J. S.
- and Bernath, P.: IASI carbon monoxide validation over the Arctic during POLARCAT spring
- 783 and summer campaigns, Atmos. Chem. Phys., 10(21), 10655–10678, doi:10.5194/acp-10-
- 784 10655-2010, 2010.
- Pope, A., Burnett, R., Thun, M., EE, C., D, K., I, K. and GD, T.: Long-term Exposure to Fine
- 786 Particulate Air Pollution, J. Am. Med. Assoc., 287(9), 1132–1141,
- 787 doi:10.1001/jama.287.9.1132, 2002.
- 788 Raleigh, C., Linke, A., Hegre, H. and Karlsen, J.: Introducing ACLED: An Armed Conflict
- Location and Event Dataset: Special Data Feature, J. Peace Res., 47(5), 651–660,
- 790 doi:https://doi.org/10.1177/0022343310378914, 2010.
- Ramo, R., Roteta, E., Bistinas, I., van Wees, D., Bastarrika, A., Chuvieco, E. and van der Werf,
- 792 G. R.: African burned area and fire carbon emissions are strongly impacted by small fires
- ⁷⁹³ undetected by coarse resolution satellite data, Proc. Natl. Acad. Sci. U. S. A., 118(9), 1–7,
- 794 doi:10.1073/pnas.2011160118, 2021.
- Robinson, T. P., Wint, G. R. W., Conchedda, G., Van Boeckel, T. P., Ercoli, V., Palamara, E.,
- 796 Cinardi, G., D'Aietti, L., Hay, S. I. and Gilbert, M.: Mapping the Global Distribution of

- ⁷⁹⁷ Livestock, PLoS One, 9(5), e96084, doi:10.1371/journal.pone.0096084, 2014.
- 798 Roteta, E., Bastarrika, A., Padilla, M., Storm, T. and Chuvieco, E.: Development of a Sentinel-2
- 799 burned area algorithm: Generation of a small fire database for sub-Saharan Africa, Remote
- 800 Sens. Environ., 222(November 2018), 1–17, doi:10.1016/j.rse.2018.12.011, 2019.
- 801 Rufino, M. C., Rowe, E. C., Delve, R. J. and Giller, K. E.: Nitrogen cycling efficiencies through
- 802 resource-poor African crop-livestock systems, Agric. Ecosyst. Environ., 112(4), 261–282,
- 803 doi:10.1016/j.agee.2005.08.028, 2006.
- 804 Sauvage, B., Gheusi, F., Thouret, V., Cammas, J. P., Duron, J., Escobar, J., Mari, C., Mascart, P.
- 805 and Pont, V.: Medium-range mid-tropospheric transport of ozone and precursors over
- Africa: Two numerical case studies in dry and wet seasons, Atmos. Chem. Phys., 7(20),
- 807 5357–5370, doi:10.5194/acp-7-5357-2007, 2007.
- 808 Stevens, C. J., David, T. I. and Storkey, J.: Atmospheric nitrogen deposition in terrestrial
- 809 ecosystems: Its impact on plant communities and consequences across trophic levels,
- 810 Funct. Ecol., 32(7), 1757–1769, doi:10.1111/1365-2435.13063, 2018.
- 811 Sutton, M. A., Reis, S., Riddick, S. N., Dragosits, U., Nemitz, E., Theobald, M. R., Tang, Y. S.,
- 812 Braban, C. F., Vieno, M., Dore, A. J., Mitchell, R. F., Wanless, S., Daunt, F., Fowler, D., Blackall,
- 813 T. D., Milford, C., Flechard, C. R., Loubet, B., Massad, R., Cellier, P., Personne, E., Coheur, P. F.,
- 814 Clarisse, L., Van Damme, M., Ngadi, Y., Clerbaux, C., Skøth, C. A., Geels, C., Hertel, O., Wichink
- 815 Kruit, R. J., Pinder, R. W., Bash, J. O., Walker, J. T., Simpson, D., Horváth, L., Misselbrook, T. H.,
- 816 Bleeker, A., Dentener, F. and de Vries, W.: Towards a climate-dependent paradigm of
- ammonia emission and deposition, Philos. Trans. R. Soc. B Biol. Sci., 368(1621), 20130166-
- 818 20130166, doi:10.1098/rstb.2013.0166, 2013.
- 819 Tian, D. and Niu, S.: A global analysis of soil acidification caused by nitrogen addition,
- 820 Environ. Res. Lett., 10(2), 024019, doi:10.1088/1748-9326/10/2/024019, 2015.
- 821 USDA Agricultural Air Quality Task Force: Ammonia Emissions : What To Know Before You
- 822 Regulate, Washington, DC. [online] Available from:
- http://www.nrcs.usda.gov/wps/portal/nrcs/detail/national/air/taskforce/?cid=stelprdb1
 268645, 2014.
- 825 Vitousek, P., Naylor, R., Crews, T., David, M., Drinkwater, L., Holland, E., Johnes, P.,
- 826 Katzenberger, J., Martinelli, L. A., Matson, P. A., Nziguheba, G., Ojima, D., Palm, C. A.,
- 827 Robertson, G., Sanchez, P., Townsend, A. and Zhang, F.: Nutrient Imbalances in Agricultural

- 828 Development, Science (80-.)., 324, 1519–1520, 2009.
- 829 Di Vittorio, C. A. and Georgakakos, A. P.: Land cover classification and wetland inundation
- 830 mapping using MODIS, Remote Sens. Environ., 204(May 2017), 1–17,
- 831 doi:10.1016/j.rse.2017.11.001, 2018.
- 832 Vrieling, A., de Beurs, K. M. and Brown, M. E.: Variability of African farming systems from
- 833 phenological analysis of NDVI time series, Clim. Change, 109(3–4), 455–477,
- 834 doi:10.1007/s10584-011-0049-1, 2011.
- 835 De Wachter, E., Barret, B., Le Flochmoën, E., Pavelin, E., Matricardi, M., Clerbaux, C., Hadji-
- 836 Lazaro, J., George, M., Hurtmans, D., Coheur, P. F., Nedelec, P. and Cammas, J. P.: Retrieval of
- 837 MetOp-A/IASI CO profiles and validation with MOZAIC data, Atmos. Meas. Tech., 5(11),
- 838 2843–2857, doi:10.5194/amt-5-2843-2012, 2012.
- 839 Warner, J. X., Dickerson, R. R., Wei, Z., Strow, L. L., Wang, Y. and Liang, Q.: Increased
- 840 atmospheric ammonia over the world's major agricultural areas detected from space,
- 841 Geophys. Res. Lett., 44(6), 2875–2884, doi:10.1002/2016GL072305, 2017.
- 842 Whitburn, S., Van Damme, M., Kaiser, J. W., Van Der Werf, G. R., Turquety, S., Hurtmans, D.,
- 843 Clarisse, L., Clerbaux, C. and Coheur, P. F.: Ammonia emissions in tropical biomass burning
- 844 regions: Comparison between satellite-derived emissions and bottom-up fire inventories,
- 845 Atmos. Environ., 121, 42–54, doi:10.1016/j.atmosenv.2015.03.015, 2015.
- 846 World Bank: World Bank Open Data, World Bank Open Data [online] Available from:
- 847 https://www.data.worldbank.org (Accessed 2 February 2019)
- 848 Yegbemey, R. N., Kabir, H., Awoye, O. H. R., Yabi, J. A. and Paraïso, A. A.: Managing the
- 849 agricultural calendar as coping mechanism to climate variability: A case study of maize
- 850 farming in northern Benin, West Africa, Clim. Risk Manag., 3, 13–23,
- doi:10.1016/j.crm.2014.04.001, 2014.
- 852 Yokelson, R. J., Christian, T. J., Karl, T. G. and Guenther, A.: The tropical forest and fire
- 853 emissions experiment: laboratory fire measurements and synthesis of campaign, Rev. Int.
- Acupunt., 8, 3509–3527, doi:10.1016/s1887-8369(09)71579-0, 2008.
- 855 Zheng, B., Chevallier, F., Ciais, P., Yin, Y. and Wang, Y.: On the Role of the Flaming to
- 856 Smoldering Transition in the Seasonal Cycle of African Fire Emissions, Geophys. Res. Lett.,
- 45(21), 11,998-12,007, doi:10.1029/2018GL079092, 2018.
- 858