

Reply to reviewers, “Radiative and chemical implications of the size and composition of aerosol particles in the existing or modified global stratosphere” by Daniel M. Murphy et al.

We thank the reviewers for their detailed comments that have improved the manuscript.

The comments are numerous so we are putting replies in-line with the comments. Since some formatting is lost when posted, our replies start with a dash in order to distinguish them from the reviewers’ comments.

## **Anonymous Referee #1**

Received and published: 22 October 2020

The manuscript “Radiative and chemical implications of the size and composition of aerosol particles in the existing or modified global stratosphere” by Murphy et al., describes measurements of stratospheric size resolved aerosol composition from optical particle counters and the PALMS aerosol mass spectrometer flown on the ATom aircraft missions in 2016-2018. The authors discuss the compositional signatures that could be indicative of aerosol transport and formation pathways that yield aerosol with two distinctly different size modes. The radiative implications of these two distinct size modes are discussed and finally the implications of these for future geoengineering. Significant anomalies in stratospheric aerosol loading and composition that are associated with volcanic eruptions, stratospheric smoke injections and transport of dust into the stratosphere are mentioned but not discussed in detail. Efforts to simulate the observed stratospheric aerosol size distributions using the CESM/CARMA model are briefly described.

The measurements of size resolved composition of lower stratospheric aerosol described in this work are unique, novel and of significant scientific importance. As this manuscript correctly describes, these results have broad implications for stratospheric chemistry and radiation and thus climate. The authors present an overview of the composition measurements in sections 1-4 at a high level that provides a brief description of the inherently complex topic of size dependent aerosol composition in the UTLS. However, the superficial description and analysis of the data set presented in this manuscript is insufficient to support the broad and generalized observations and conclusions presented in Sections 5 and 6. Many of these observations and conclusion are likely to be correct, and may be supported by this ATom data set, but have not been demonstrated to the reader in this work.

- We appreciate the sentiment that the measurements are “unique, novel, and of significant scientific importance.”

As this appears to be the first description of the PALMS composition measurements from the ATom campaign, the brief and qualitative description of the sampling and do not allow for the reader to understand how representative or significant the compositional analysis is, how definitively tropospheric and stratospheric air masses are separated, or even basic information such as how the ozone measurements were acquired. While the authors state that about 10,000 particles were sampled in the Northern Hemisphere (line 238) there is no indication of how these particles were divided between the various flights, how many flights there were, at what locations these particles were sampled, what fraction of the data was from tropopause folds, or how many particles were measured in the Southern Hemisphere. Figure S1 provides some insight into where the measurements may have been made, but as the color bar in S1 does not correspond with the 250 ppbV cut off for stratospheric air, even this provides insufficient information. Due to the low ceiling of the DC-8, the authors indicate that the sampled “stratospheric air was always associated with low tropopauses, sometimes in tropopause folds”. Given that the upper tropospheric organic aerosol loading can be very substantial, particularly with the wild fire activity likely present for ATom1 and ATom3, it is important to know whether these data are from stratospheric measurements above the broader tropopause (less likely in the summer time) or from tropopause folds (more likely in the summer time) and how edge effects and mixing were considered for the latter.

- We have added Table S.1 of when and where the DC8 was in the stratosphere and the relevant number of mass spectra acquired. A citation has been added for the ozone instrument. The way edge effects and mixing for tropopause folds was considered is largely through the use of ozone as a criterion: excessive mixing with tropospheric air will reduce the ozone levels below 250

ppbv. This is now more clearly stated. We have added tropopauses to Figure S1; they show that most of the data with more than 250 ppbv of ozone in all seasons is above a high latitude tropopause. We have changed the color scale on Figure S1 that starts at 250 ppbv.

Without the presentation of a more thorough analysis of these data, it is not clear that some of the major conclusions of the paper are solidly supported. It may be the case that the transport of tropospheric organic-sulfate aerosols is producing a second smaller stratospheric aerosol mode centered around 200 nm in some specific cases. But this is not robustly supported in 5 of the 8 cases shown in Figure 1. In ATom4 NH, the peak near 200 nm appears to be primarily driven by an increase in pure sulfuric acid aerosol. In only one of the SH cases (ATom 3) is there convincing evidence of a secondary mode, and again in this case it appears to be driven by sulfuric acid aerosol (presumably from Calbuco). Without much more detailed description and analysis it is hard to be sure that the conclusion that tropospheric aerosol significant to the global lower stratosphere (paragraph starting on line 450) are really supported by this data set. This would be a very important result and I encourage the authors to layout a more convincing case based on the ATom data set.

- The referee's comment about not seeing that tropospheric particles are smaller shows that the size distributions weren't all that visible in the stacked plots of Figures 1 and 3. They also needed additional explanation. We've added a new figure (4) and associated text that show this more clearly. We have also added text (near new line 230) saying that the NH ATom4 case noted by the reviewer was in a dust event with strong transport from near-surface sources of particle other than dust. We appreciate the reviewer pointing out something that may not have been obvious to readers.

In section 4, the authors state "As expected for their sources, the concentration of meteoric-sulfuric particles increases with altitude and the concentration of organic- sulfate particles decreases with altitude." While it does appear to be clear that concentration of meteoric-sulfuric aerosol increases with ozone/altitude, the data presented here do not appear to provide convincing support for a general decrease in organic- sulfate particles with altitude. This relationship appears to be fundamental to apportioning the source of these particles to transport across the tropopause. Three of the organic-sulfate profiles shown in Figure 4 appear to be constant with respect to ozone/altitude, three may show a decrease and one shows an increase. The significance of these slopes is hard to determine without some metric of the measurement uncertainty. From the narrative, it appears that there was a limited (but unspecified) amount of particle composition data at ozone levels above 500 ppbv, so an indication of how significant the measurements > 500 ppbv is necessary to interpret the importance of this slope. Comparing the ATom 2 data in Fig 3 with Fig 1, it appears that there is in fact no significant difference in organic-sulfate particles in the 200nm mode, and actually a significant increase in the amount of organic-sulfate in the 600nm mode. The tropospheric origin of these organic-sulfate particles could also be indicated by their compositional signature, as the authors state on line 143. This is not demonstrated to the reader, even though the data must surely be available.

- It is important to understand that the identification of the tropospheric particles is based on their mass spectra, not the vertical profile. We have changed the text in several places to state this more clearly. We have added a new supplemental figure S2 showing how closely the mass spectra of organic-sulfate particles in the stratosphere match those in the upper troposphere. (We have this in supplemental because mass spectra of the major types of particles in the lower stratosphere have already been published in Murphy et al., 2014.) The text in question has also been revised to point out that ATom3, which has increasing concentrations with ozone, is a special case because of the pyrocumulus injection. The other cases show decreasing concentrations whenever the concentrations are significant. The new table S1 shows that there are many mass spectra of particles in air with more than 500 ppbv ozone for the deployments when the DC8 sampled such air.

Throughout the paper there is no quantitative discussion of the uncertainty or statistics associated with this highly averaged data (in the narrative or in any of the figures). This issue is addressed anecdotally when convenient. For example, in the caption in figure 1 the authors do state that there 'may be artifacts' due to Mie resonance for certain

features but other features are robust without any further explanation or justification. If there were other publications describing this data set in detail, some of these considerations could be addressed through references to these publications, however it appears that these publications are planned for the future.

- We have added a paragraph on uncertainties. The size measurements are described in much more detail in the Brock et al. (2019) reference, which is specifically about the ATom measurements.

Section 5 of the manuscript primarily describes the radiative and chemical impact that arises from the divergent size modes between aerosol of tropospheric and stratospheric origin. The narrative in section 5 is somewhat disjointed and difficult to follow as it is not always clear how it relates back to the measurements presented in the first sections of the paper. For instance, in section 5.1, the infrared absorption is only discussed in the context of particle volume, yet while the authors acknowledge that particle composition plays a role (line 266) this is not tied back to one of the primary observational findings, that the two size modes have significantly different composition. There are also several results that described in section 5 that are not demonstrated to the reader. For example, on line 329: "The relative contributions to scattering of light at  $< 240$  nm are fairly similar to the surface area panel in Figure 7 except that sizes smaller about 80 nm and larger about 600nm contribute less to UV scattering than they do to surface area." While this statement is quite likely true, it has not been shown and it is hard to quantify what 'fairly similar' describes. A similar statement begins on line 313 "On Figure 7, the sedimentation flux as a function of size would be slightly more skewed to large diameters than the light scattering panel." Such generalized statements need to be shown to be supported by the data.

- The text in section 5.1 has been changed to clarify that while the amount of infrared absorption may change with composition, the size dependence does not. We have added a new figure to the supplemental material with the additional curves about sedimentation and UV scattering. We are trying in the main text not to have figures with too many curves.

Section 6 of the manuscript sets out to describe the relevance of this work to volcanic eruptions or to future geoengineering projects. While the discussion is interesting it is highly speculative and largely unrelated to the data presented in Sections 1-4. For example, in the paragraph starting on line 396 it is implied that differences between the modeled sulfuric acid particle sizes and these measurements (Figure 2) is a consequence of one of the main results that is presumably shown by the data – multiple sources of stratospheric aerosol. However, no details on the initialization for the model run or analysis of what is driving these discrepancies is provided, as a result the comparison in Figure 2 is largely anecdotal.

- The model is described in more detail in the Yu et al. reference and we have added a sentence saying that the model does not distinguish meteoric-sulfuric particles.

Finally, the title of the paper "Radiative and chemical implications of the size and composition of aerosol particles in the existing or modified global stratosphere" is misleading. While the aerosol composition data is presented in a limited way, the implications (described in sections 5 and 6) are based entirely on the size distribution of the particles not their composition, even when composition would certainly be important to these implications (e.g. 5.2 Scattering and 5.3 chemistry).

- We feel the title is appropriate. The paper does discuss both size and composition. Although the composition does not enter directly into the radiative properties, it is crucial to the attribution of those properties. Without information about the composition we could not say, for example, what percentage of extinction is due to particles formed in the stratosphere and what percentage due to particles formed in the troposphere. We have added two sentences about this near the beginning of section 6.

Technical Comments:

Line 54: “The moles of oxygen were approximately four times sulfur plus about 0.2 times carbon” - rephrase to make it clear where the brackets are  $O_2 = 4 * (S + 0.2C)$ ?

- This has been rephrased and a comma added to make it more clear where the brackets belong.

Paragraph starting line 95: It is unclear why the OPC data is needed if PALMs provides sizing?

- More explanation has been added.

Line 174: Using ozone as a proxy for altitude is understandable, but some indication of what an equivalent tropopause relative altitude or range of altitudes that a given ozone mixing ratio corresponds with would be useful.

- This is shown in Figure S1.

Figure 6: What is the unit of measure for ‘Approximate net cooling’ and how was this calculated?

- The scaling is consistent with the scattering curves. For example, if the “net cooling” curve is 80% that of the light scattered to outer space curve, this means that the cooling of the Earth is 80% of what it would be if only light scattering were considered. Text has been added to section 6.2 stating this. Section 6.2 also describes the calculation.

Figure 7: The units of these plots are unclear. Either provide units, or normalize the data to make it clear that it is relative surface/scatter/volume

- Units have been added to the figure.

Figure 8: This figure is confusing. What is meant by the IR-heating trace? Is the peak at 0.5 the most relative cooling per IR heating, and would that mean the least or most IR heating?

- More explanation has been added to the caption. The reviewer is correct, the peak at 0.5 means the most cooling per IR heating.

Figure S1: It would be useful to start the color bar at 250ppbv so that only the points used in this study are highlighted, with a clear distinction for where the particles shown in Figure 3 were found.

- We have changed the color scale on Figure S1 that starts at 250 ppbv. Adding a separate color and symbol size for the 100-250 ppbv range of ozone allows readers to see near-tropopause locations. We’ve also chosen a color scale that doesn’t rely on green-red distinctions.

## Anonymous Referee #2

Received and published: 4 November 2020

The LMS is an important part of the stratosphere and difficult to observe. Its composition differs from the rest of the stratosphere in many ways due to the mixing of tropospheric and stratospheric air. The study by Murphy et al. is based on in-situ data from aircraft of aerosol particle size distributions and composition of individual particles. Particles were classified depending on their composition to study the history of the particles. Radiative impact was discussed in relation to volcanic eruptions and climate engineering.



I find large issues with the authors' interpretation of the data. One reason is the generalization in sections 5-7 based on few data (a single season in a single year). Another is the data on organics, and the claim that the organic aerosol comes from the troposphere only, which contradicts previous work. No uncertainties are presented for the observations as far as I can tell, and there are no statistical analysis to support the claims of trends.

- The reviewer is incorrect that sections 5 to 7 are based on a single season or year. The calculations are not dependent on the season but represent the general consequences of most tropospheric particles being smaller than most stratospheric particles. We picked one season to serve as an example rather than have multiple panels for every figure. Summary data for the other seasons and hemispheres are shown in Figure S4.

I find several shortcomings in the manuscript and cannot recommend publication without changes in the data analysis / interpretation of the data.

#### Major comments

1. Uncertainties and quantification a. How large are the uncertainties in the observations? I don't find any numbers on that. b. How well can one quantify different aerosol constituents from the PALMS data? c. Is it possible to tell whether there is a trend, as in Figure 4, if there are no statistical analysis to support it?

- We have added a paragraph on uncertainties.

2. There are too little organics in the data compared to previous work, and this is not discussed in relation to those studies. a. There is no discussion on organics after Calbuco. Several other groups have found large amounts of organics in volcanic aerosol. Two examples are Schmale 2010 and Andersson 2013 reported much higher organics (or carbon) to sulfate ratios in volcanic particles. Those data were from AMS and Ion beam analysis. b. Vertical gradient in organics is different from Martinsson 2019, who reported higher carbon abundance deeper into the LMS. I cite from their abstract: "...the carbonaceous and sulfurous components of the aerosol in the lowermost stratosphere (LMS) show strong increases in concentration connected with springtime subsidence from overlying stratospheric layers. The LMS concentrations significantly exceed those in the troposphere, thus clearly indicating a stratospheric production of not only the well-established sulfurous aerosol, but also a considerable but less understood carbonaceous component. ..." c. There is no real discussion on wildfire smoke. The smoke from the Aug 2017 fires in western North America is evident in Figure 4. The relatively small impact on the organics after the wildfire in that figure is strange given that the event was almost volcanic sized (Peterson 2018). In Figure 1, it is apparently less organics 1-2 months after the fire than in spring the next year. Why is that? The discrepancy between the data in the manuscript with data from other studies is not discussed as far as I can tell. Identifying particles containing organics is not the same as measuring the mass (of organics or carbon), which other techniques do.

- There are few measurements of organic material in the stratosphere other than from PALMS. Our data seem consistent with the CARIBIC bulk analyses of Nguyen and Martinsson (2007).

- The literature consensus is that volcanic aerosol (excluding ash similar particles immediately after an eruption) is sulfuric acid with little organic content. We have added two references. The papers cited by the reviewer did not definitively find higher organics or carbon in volcanic particles. Schmale et al. 2010 state that their data are only tentative: "Overall, the quantification of organic material in the volcanic plume is subject to uncertainties so that we cannot state whether there is a true increase in organics". In addition, both Schmale et al. and Andersson et al., 2013 ascribe tropospheric particles as a likely cause of organic aerosol in a volcanic plume. This is more consistent with our analysis than the reviewer suggests. To quote Schmale et al. "It is unclear whether the apparent increase in carbonaceous mass might reflect injection of volcanic species or injection of tropospheric species which experienced entrainment into the eruption column." And from Andersson et al "we hypothesize that organic material in entrained air

constitutes a significant fraction of the particulate carbon observed in volcanic clouds.” If entrainment does occur, the PALMS single particle analysis would identify tropospheric particles entrained into a volcanic plume as tropospheric rather than volcanic.

- The inferences in Martinsson et al. 2019 are based on bulk analysis without size resolution. We think it is clear from this manuscript that at least size-resolved data, if not single particle data, are crucial to understanding particle sources in the lower stratosphere. Previous PALMS data support this, see Murphy et al. (2007). It is important to distinguish the organic content of the overall aerosol, which decreases above the tropopause, with the organic content when considering only the tropospheric particles, which has little vertical gradient. This is shown in the figures in Murphy et al. (2007) and also Figure 6 (was Figure 5) in the current manuscript.

3. Large part of the manuscript is focused on a single season in the Northern Hemisphere (Atom2). It is unclear to me why this is the case. Sulfate concentrations in the LMS varies with season due to both seasonal variation in subsidence from the stratospheric overworld and varying cross TP transport. Stratospheric influence is large in winter and spring, and low in summer and fall. Thus: a. Data presented for a specific season is representative for that season only, and not for the entire year. b. General climatic conclusions cannot be drawn from a single season in a single year.

- The reviewer is incorrect that the analysis is based on a single season or year. The calculations are not dependent on the season but represent the general consequences of most tropospheric particles being smaller than most stratospheric particles. We picked one season to serve as an example rather than have multiple panels for every figure. Summary data for the other seasons and hemispheres are shown in Figure S4.

4. Data after the Calbuco eruption are almost not discussed at all. The authors mention the eruption and that it had some impact on the sulfuric acid but no more details. I understand that the authors focused on the Northern Hemisphere, but this omission is strange to me. I expect some discussion on data that are included in a manuscript.

- We have added additional discussion of the Calbuco data, especially implications for remote retrievals.

#### Other comments

1. I would like to see a more clear discussion on the history of sulfuric particles. Meteoric ones from the upper stratosphere, and pure sulfuric from the lowest stratosphere (directly from the tropical lower stratosphere in the BDC).

- There is more discussion of this in a previous paper (Murphy et al., 2014) and we hope to continue to acquire data at higher altitudes that will help further study of the history of the types of sulfuric acid particles.

2. The ExTL has very different composition than the rest of the LMS and (stratosphere). The author never mention the ExTL. Why is that?

- The ExTL concept turns out not to be very helpful in presenting these particular data.

3. I think that the phrase biomass burning shall be changed to wildfire smoke since it comes from PyroCb intrusion(s) to the stratosphere. Biomass burning leads the reader to believe that it is a general upwelling from diffuse fires instead of PyroCb formation in enormous fires.

- The use of “biomass burning” to refer to smoke particle in the remote atmosphere follows common usage in most of the literature. We have added the modifier “wildfire” where appropriate.

4. I think that the manuscript should have a concise conclusions section after the C3 discussions section.

- For this manuscript we found that the text flowed better if the discussion and conclusions were combined. We have added a short paragraph that summarizes some of the implications for satellite retrievals.

L45. “...The local tropopause and slightly above the altitude of the tropical tropopause...”. This is not true. The tropical TP is located at ~17 km, and the LMS ex- tends to ~14-15 km in the extratropics. L135. “...The primary source of sulfuric acid in the stratosphere, oxidation of carbonyl sulfide, is similar in the two hemispheres...”. Do you mean the background stratospheric aerosol? It is still an open question whether SO<sub>2</sub> or COS is the most important for the background stratospheric aerosol. L183. Change to: ...Northern hemisphere than in the Southern Hemisphere... L191. I do not understand the statement: “...the organic-sulfate particles contain on average about 40 to 80 organic material by mass...” L260. “...Alumina with the size distribution from rocket emissions was calculated to cause net warming (Ross and Sheaffer, 2014)...”. This is a strange comment in the end of a section on volcanic particles and IR. Are there a sentence missing? L361. Change “will it” to “it will” L469. “...the modeled particles have about 65% of the climate impact per unit mass as the observations, 160% of the surface area, and sediment about 60% as fast...” Is this referenced to findings in the manuscript? Are there actual observations of these quantities or do you mean estimated from the PALMS observations?

- Minor text changes suggested here have been made. The definition of the lowermost stratosphere at follows the references listed. There is actually a lot of spread in the definition of “lowermost stratosphere” by different authors. The question of SO<sub>2</sub> or COS as the most important source of sulfuric acid in the background stratosphere has been largely settled (Kremser et al. 2016 and Rollins et al. 2017 references). To quote the review paper by Kremser et al. “*OCS makes the largest contribution to the aerosol layer*” [italics in original]. The statement about organic-sulfate particles was missing “percent”. The text has been modified to better tie the Ross and Sheaffer statement to the text around it. The text has been changed to say that the 65% etc. numbers are compared to those calculated from PALMS observations.

- Additional changes not requested by the reviewers:

- Fine-tuning some data processing has resulted in insignificant changes in the figures and tables. For example, the volumes of meteoric-sulfuric particles and sulfuric particles in the top row of Table 1 have changed from 0.045 and 0.109 to 0.046 and 0.103 respectively. Figure 5 has been revised by averaging over narrower ranges of ozone near the bottom of the profiles where there are many mass spectra and wider range of ozone near the tops of the profile where there are fewer mass spectra. This improves the statistical weighting but the patterns of all of the curves are very similar to the previous version. The revised manuscript also corrects an error in the Figure 5 label which should have been volume instead of mass.

# Radiative and chemical implications of the size and composition of aerosol particles in the existing or modified global stratosphere

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**Abstract.** The size of aerosol particles has fundamental effects on their chemistry and radiative effects. We explore those effects using aerosol size and composition data in the lowermost stratosphere along with calculations of light scattering. In the size range between about 0.1 and 1.0  $\mu\text{m}$  diameter (accumulation mode), there are at least two modes of particles in the lowermost stratosphere. The larger mode consists mostly of particles produced in the stratosphere and the smaller mode consists mostly of particles transported from the troposphere. The stratospheric mode is similar in the Northern and Southern hemispheres whereas the tropospheric mode is much more abundant in the Northern Hemisphere. The purity of sulfuric acid particles in the stratospheric mode shows that there is limited production of secondary organic aerosol in the stratosphere, especially in the Southern Hemisphere. Out of eight sets of flights sampling the lowermost stratosphere (four seasons and two hemispheres) there were three with large injections of specific materials: volcanic, biomass burning, or dust. The ~~two~~ stratospheric and tropospheric modes have very different roles for radiative effects on climate and for heterogeneous chemistry. Because the larger particles are more efficient at scattering light, most of the radiative effect in the lowermost stratosphere is due to stratospheric particles. In contrast, the tropospheric particles can have more surface area, at least in the Northern Hemisphere. The surface area of tropospheric particles could have significant implications for heterogeneous chemistry because these particles, which are partially neutralized and contain organics, do not correspond to the surfaces used for laboratory studies of stratospheric heterogeneous chemistry. ~~The purity of sulfuric acid particles in the stratospheric mode shows that there is limited production of secondary organic aerosol in the stratosphere, especially in the Southern Hemisphere. Out of eight sets of flights sampling the lowermost stratosphere (four seasons and two hemispheres) there were three with large injections of specific materials: volcanic, biomass burning, or dust.~~ We then extend the analysis of size-dependent properties to particles considered for intentional climate modification. There is no single size that will simultaneously optimize the climate impact relative to the injected mass, infrared heating, potential for heterogeneous chemistry, and undesired changes in direct sunlight. In addition, light absorption in the far ultraviolet is identified as an issue requiring more study for both the existing and potentially modified stratosphere.

## 1 Introduction

Stratospheric particles have been studied for over 50 years (Junge and Manson, 1961; reviewed by Kremser et al., 2016). The stratospheric aerosol layer has a maximum in mixing ratio between about 20 and 25 km altitude. However, the larger air density at lower altitudes means that the majority of the mass of stratospheric aerosol is in the lowermost stratosphere, below the maximum in mixing ratio (Yu et al., 2016).

Various trends have been reported for the background stratospheric aerosol at times not influenced by major volcanic eruptions. Deshler et al. (2006) concluded there was little long-term change in background stratospheric aerosol from 1971 to 2004. Hofmann et al. (2009) found an increasing trend after 2000, and Friberg et al. (2014) found an increasing trend from 1999 to 2008. There has been recognition that moderate volcanic eruptions frequently influence the stratospheric aerosol and true nonvolcanic “background” concentrations are not necessarily present just because there has been no Pinatubo-scale eruption (Solomon et al., 1996, Vernier et al. 2011). Different altitudes may exhibit different trends (Khaykin et al., 2017). Moderate volcanic eruptions tend to mask trends in the non-volcanic background (Kremser et al., 2016). The data shown here reinforce the notion of modest but frequent perturbations to the lower stratosphere.

The overall circulation of air in the stratosphere, with rising air in the tropics and descending air in the extratropics, is mostly fed by air entering at the tropical tropopause. The lowermost stratosphere is a region at middle and high latitudes between the local tropopause and slightly above the altitude of the tropical tropopause. Air in this lowermost stratosphere is affected by both downward motion in the extratropical stratosphere and adiabatic mixing with the troposphere (Holton et al., 1995). The tropospheric influence can extend as high as about 450 K potential temperature (Rosenlof et al., 1997). All of the data described here are in the lowermost stratosphere and show the influence of both air from higher in the stratosphere and air from the troposphere.

The chemical composition of particles in the lower stratosphere has been measured by several techniques. Impactor samples collected from the CARIBIC platform have been quantitatively analyzed for N, O, S, K, Fe, and other elements (Nguyen and Martinsson, 2007; Friberg et al, 2014; [Martinsson et al., 2019](#)). The moles of oxygen were approximately four times [the moles of sulfur](#), plus about 0.2 times [the moles of carbon](#), indicating SO<sub>4</sub> in sulfate and sulfuric acid plus some contribution from oxygenated organics. Much of the detailed information on the chemical composition of aerosols in the lower stratosphere has come from the Particle Analysis by Laser Mass Spectrometry (PALMS) instrument (Murphy et al., 1998). These data show that most particles larger than about 110 nm fall into three distinct types: sulfuric acid with or without meteoric metals and mixed organic-sulfate particles from the troposphere (Murphy et al., 2014). Because the pure sulfuric acid particles do not contain biomass burning residues, they are not simply tropospheric particles that have lost organics after entering the

65 stratosphere (Murphy et al., 2007). Recent data from another single particle mass spectrometer have found comparable abundances of sulfuric acid particles with meteoric metals (Schneider et al., in discussion).

We extend the previous results to show that the mixed organic-sulfate particles from the troposphere are generally smaller than both types of sulfuric acid particles. We then show how this size difference has significant implications for light scattering and heterogeneous chemistry.

## 2 Methods

This paper includes data in the lowermost stratosphere from the Atmospheric Tomography (ATom) mission with deployments in four seasons during 2016 to 2018. Although the ATom mission was not specifically designed to sample the stratosphere, it encountered stratospheric air in both the Northern and Southern Hemispheres during its regular vertical profiles. Stratospheric air was encountered periodically at altitudes greater than about 8-7 km and latitudes poleward of ~~30~~about 30 degrees north or ~~40-35~~ degrees south (Figure S1 and Table S1). Because the NASA DC8 aircraft has a ceiling of about 12 km, stratospheric air was always associated with low tropopauses, sometimes in tropopause folds. We therefore use ozone rather than altitude as the primary definition of how far into the stratosphere measurements were taken. If there were mixing with a great deal of tropospheric air it would reduce the ozone concentration below the thresholds we set for stratospheric measurements. Any high ozone Data below 7 km altitude was were excluded just in case there were unusually high ozone concentrations at low altitude.

Size distributions for accumulation-mode particles were measured using two modified commercial laser optical particle spectrometers, an ultra-high sensitivity aerosol spectrometer (UHSAS; Droplet Measurement Technologies, Longmont, USA) from 0.07 - 0.6  $\mu\text{m}$  diameter and a laser aerosol spectrometer (LAS, TSI Inc., St. Paul, USA) from 0.6 to  $\sim 4.8 \mu\text{m}$  diameter (Kupc et al., 2018; Brock et al., 2019). The diameters are based on calibration by ammonium sulfate particles. The size resolution of the reported data is 20 bins per decade of particle size. Data are recorded at 1s intervals although averaging is needed in the stratosphere to improve counting statistics for particles in the LAS size range (Brock et al., 2019). Ozone measurements are described in Bourgeois et al., 2020.

Particle composition was measured with the Particle Analysis by Laser Mass Spectrometry (PALMS) instrument (Thomson et al., 2000; Froyd et al. 2019). A pressure-controlled aerodynamic focusing inlet brings particles into a vacuum where they cross two continuous laser beams. The transit time between the beams measures the aerodynamic diameter of each particle. The aerodynamic diameters are under vacuum conditions with most particles much smaller than the mean free path at the inlet exhaust. ~~T; however~~ transition flow corrections are considered. Transit times were calibrated to known particle sizes before and after every field deployment. A 193 nm pulse from an excimer laser is triggered when a particle arrives at the second laser beam. Either positive or negative ions are analyzed with a time-of-flight mass spectrometer. The polarity was switched every

few minutes. Most of the data shown here are from positive ion spectra. Negative ion spectra do not distinguish sulfuric acid with and without meteoric metals because the metal ions only appear in the positive ion spectra.

The optical size distributions are combined with the PALMS single-particle composition data for particles larger than 100 nm to create size distributions that are resolved by composition. In each size range, the number of particles is obtained from the optical size distributions and the fraction of particles with different compositions is obtained from PALMS. This combination requires converting the PALMS aerodynamic diameters to correspond to the optical diameters (Froyd et al., 2019). The composition-resolved size distributions presented here use wider size bins than the native optical particle counter resolution but narrower size bins than the standard AToM products (Froyd et al., 2019). The narrower bins are possible because of improved statistics after averaging over all of the data within a specific band of ozone and latitude, even if those data were not contiguous in time.

For the purpose of this study particles are classified into four basic categories: sulfuric acid with and without meteoric metals, mixed organic-sulfate particles, and other particles including dust. Example spectra are shown in Murphy et al. (2014). When appropriate further distinctions can be made, such as separating biomass burning particles from other mixed organic-sulfate particles.

Sources of uncertainty include particle statistics, identifying particle types, the particle volume from the optical particle counters, and uncertainty in combining the PALMS spectra with the optical size distributions. Overall statistics are excellent, with approximately 800,000 single particle mass spectra acquired above 7 km altitude during AToM. Of those, approximately 78,000 were positive ion spectra at ozone concentrations larger than 250 ppbv (Table S1). However, low ambient number densities for large particles mean that there are few mass spectra of particles larger than about 1  $\mu\text{m}$ , leading to statistical noise for large particles visible in the figures in this paper. Except for AToM4 Northern Hemisphere, particles larger than 1  $\mu\text{m}$  contributed only a very small fraction of the aerosol volume in the lower stratosphere. Classification uncertainty depends on the type of particle. The meteoric-sulfuric particles are very distinctive. From manually examining mass spectra we estimate that only a few percent are incorrectly classified. Sulfuric acid particles are more difficult to classify because some tropospheric particles with low organic content are similar to stratospheric sulfuric acid particles. The criteria used here probably err on the side of underestimating tropospheric particles in the lower stratosphere, especially in the Southern Hemisphere. Using various criteria for separating organic-sulfate and sulfuric acid particles suggests a classification uncertainty of perhaps 25% for those particle types in the lower stratosphere. With sufficient averaging (minutes), the volume derived from optical size distributions has an uncertainty propagated from size and flow uncertainties of about +13/-28% in the accumulation mode and up to +/- 50% above 1  $\mu\text{m}$  (Kupc et al., 2018; Brock et al., 2019). Excellent agreement between extinctions calculated from the size distributions and independent extinction measurements indicates that systematic errors may actually be less than this (Brock et al., 2019). Because PALMS measures aerodynamic diameters, mapping the PALMS spectra onto the optical size

distributions requires particle density and shape (Froyd et al., 2019). These are well known for sulfuric acid particles, less so for some particles such as dust. How uncertainty in density affects the combination of PALMS and optical data depends on the shape of the size distribution and the number of types of particles. In a flat size distribution or if there is only one dominant type of particle present, uncertainty in density introduces very little additional uncertainty when combining PALMS data with size distributions. The uncertainty is larger if there are several types of particles present in a size range with rapidly changing particle volume versus size. Statistical uncertainty in the fractional organic content of mixed particles is less than 10% if at least dozens of mass spectra are averaged (Froyd et al., 2019).

Calculations of light scattered back to outer space are made for an optically thin layer uniformly spread over a sunlit hemisphere as described in Murphy (2009), except that these calculations use an atmospheric transmittance appropriate for approximately 11 km altitude (Arvesen, 1969). Changing the solar spectrum over the entire range from top-of-atmosphere to surface gives qualitatively similar results. At the low relative humidities in the stratosphere, water uptake is less important for optical properties than it is in the troposphere. The mean relative humidity for the ATom data at ozone > 250 ppbv was less than 10%.

A sectional aerosol model (CARMA) coupled with the NSF/DOE Community Earth System Model (CESM) is used in the study to simulate the composition and size distributions of stratospheric aerosols (Yu et al., 2015; Toon et al., 1988). CESM-CARMA tracks two external-mixed groups of aerosols. The first group consists of pure sulfate particles (formed through nucleation and condensation of water vapor and sulfuric acid) with 20 size bins ranging from 0.4 nm to 2.6  $\mu\text{m}$  in diameter; the second group consists of internal mixed aerosols (containing condensed sulfate, organics, black carbon, salt and dust) with 20 size bins from 0.1  $\mu\text{m}$  to 17  $\mu\text{m}$ . The model does not explicitly separate meteoric-sulfuric particles from other sulfuric acid particles. The model is run at a horizontal resolution of 1.9° (latitude) x 2.5° (longitude). It has 56 vertical layers from the surface to 1.8 hPa with a vertical resolution of ~1 km near tropopause.

### 3 Composition-resolved size distributions in the stratosphere

Figure 1 shows the composition-resolved size distributions measured in the lowermost stratosphere for the four ATom deployments, separated by the Northern and Southern Hemisphere. The data are for ozone between 250 and 400 ppbv. This range of ozone is chosen to be definitely in stratospheric air and to include data from both hemispheres on all four deployments. For comparison, the median ozone concentration at the lapse rate thermal tropopause was slightly less than 100 ppbv. On each panel the thick black line is the size distribution from the optical particle counters. At each size the fraction of particle types from PALMS is shaded. A number of features in Figure 1 are worthy of comment.

The volume distributions show a peak near 400 nm diameter and another peak, or at least a shoulder, near 180 or 200 nm. These sizes are both within what is considered the accumulation mode. The composition reveals why there are two modes.



Indeed, without the composition it would be difficult to be sure that there were two separate modes. For example, the Wyoming particle counters used on stratospheric balloon flights (Deshler et al., 2003) do not clearly resolve the modes.

170 The 400-nm mode is from sulfuric acid particles produced in the stratosphere, especially those with meteoric metals. The size of these meteoric-sulfuric particles is extremely consistent through both hemisphere and the four deployments (Table 1). The primary source of sulfuric acid in the stratosphere, oxidation of carbonyl sulfide, is similar in the two hemispheres. The meteoric-sulfuric particles also have a narrow size distribution, with a typical geometric standard deviation of about 1.4 when  
175 distributions because smaller particles have relatively more surface area for condensation. Sulfuric acid particles without meteoric material have more diverse sizes except for the volcanically influenced ATom1 Southern Hemisphere, when the sulfuric acid particles had a narrow size distribution very similar to the meteoric-sulfuric particles.

The smaller mode near 200 nm is from mixed organic-sulfate particles that have mixed into the stratosphere. The mass spectra  
180 of particles in the smaller mode are essentially identical to those of particles in the upper troposphere (Figure S2). Unlike studies that relied solely on bulk composition (e.g. Martinsson et al., 2019), we identify the tropospheric contribution based on the mass spectra of individual particles (Murphy et al., 2007). Comparing the Northern and Southern Hemispheres in Figure 1, the concentration of the smaller mode is larger in the Northern Hemisphere. The upper troposphere in the Southern Hemisphere has generally lower aerosol concentrations, so mixing in a given amount of tropospheric air will bring in fewer  
185 particles than the same amount of mixing in the Northern Hemisphere.

Of the eight cases in Figure 1, three have much higher aerosol concentrations than the others, for three very different reasons. The ATom1 Southern Hemisphere had a ~~large and very narrow~~ mode of pure sulfuric acid particles. These were most likely produced from SO<sub>2</sub> injected into the stratosphere by the Calbuco eruption in April 2015, about 16 months before the  
190 measurements (Bègue et al., 2017). The sulfuric acid particles are remarkably pure, except for associated water. This is consistent with previous data on aerosol volatility and infrared spectra after the eruption of Mt. Pinatubo (Deshler et al., 1992; Grainger et al., 1993). The size distribution is also quite narrow, with a geometric standard deviation  $\sigma_g$  of 1.29 for just the sulfuric acid particles (Table 1). Satellite aerosol retrievals may use wider size distributions. For example, Bauman et al. (2003) use a lookup range for  $\sigma_g$  of 1.3 to 2.3 and show that retrievals may not find solutions if  $\sigma_g \sim 1.1$ . The size distribution of these  
195 volcanic particles suggests caution in satellite and lidar retrievals of stratospheric aerosols dominated by one source of particles grown by condensation of sulfuric acid. As can be seen in Figure 1, overall size distributions typically are not this narrow because they are broadened with particles from multiple sources.

The ATom3 Northern Hemisphere had a large tropospheric organic-sulfate contribution. More detailed composition shows  
200 that these included a large fraction of biomass burning particles. The ATom3 flights were about two months after Canadian

fires produced a massive injection of smoke into the stratosphere and also rising through the tropopause due to diabatic heating (Torres et al., 2020). A separate paper is in preparation about the in-situ data from this wildfire event. The ATom4 Northern Hemisphere had both a large contribution from organic-sulfate particles and a remarkable amount of dust at and above the tropopause. Concentrations at the tropopause often reached several micrograms per standard cubic meter. The dust was very widespread: it was measured over both the north Atlantic and Pacific Oceans over more than 40 degrees of latitude. This may be Asian dust and pollution carried to high altitude in an event similar to that described by Huang et al. (2008). The ATom4 and Huang et al. events were both in May. A separate paper is also planned about this dust event.

Figure 2 shows the CESM/CARMA model results for the ATom2 flights for the same 250-400 ppbv range of ozone as the data in Figure 1. ATom2 is chosen because neither hemisphere was perturbed by volcanic sulfate, biomass burning, or dust. The model reproduces the tropospheric mode well in the Northern Hemisphere but overestimates it in the Southern Hemisphere. The model reproduces the total volume of stratospheric particles well in both hemispheres but the modeled diameter of these particles is too small. A possible reason is that the model does not include meteoric smoke particles on which sulfuric acid can condense. That is, the model treats both the meteoric-sulfuric and sulfuric acid particles observed by PALMS as a single type. Figure S3 compares the Northern Hemisphere ATom data to Wilson et al. (2008). Consistent with the ATom observations, the tropospheric particle mode at about 200 nm is in some cases distinct from the larger stratospheric mode.

Figure 3 shows composition-resolved size distributions further into the stratosphere with ozone between 500 and 850 ppbv. Only the Northern Hemisphere during ATom2 and ATom4 had significant amounts of PALMS data in this ozone range. For ATom2, the primary difference at higher ozone was more meteoric-sulfuric particles, a result also found by Schneider et al. (in discussion). The large mixing event of dust and other tropospheric particles during ATom2 barely affected “altitudes” of more than 500 ppbv ozone. Figure 4 has some of the data in Figures 1 and 3 replotted to emphasize the size distributions of each component. The size distributions of the meteoric-sulfuric particles are extremely consistent. The size distributions of the sulfuric acid particles without meteoric content are usually broader, possibly indicating more diverse sources. One exception is the very narrow size distribution of the particles after the Calbuco eruption. The ATom3 Northern Hemisphere has a mode of larger organic particles from the pyrocumulus injection event. The dust event in the Northern Hemisphere during ATom4 also brought up organic-sulfate and sulfuric particles, or their precursors, from near the surface as well as dust.

#### 4 Vertical profiles

The shaded regions in Figure 1 can be integrated over all sizes to give the volume associated with each type of particle, then multiplied by a density to give the mass. Figure 5 shows vertical profiles of the mass-volume concentrations for the meteoric sulfuric particles and organic-sulfate particles. As expected for their a high-altitude source, the concentration of meteoric-sulfuric particles increases with altitude ozone concentration. Except for the ATom3 wildfire biomass burning injection, the

235 concentration of organic-sulfate particles decreased with ~~altitude~~ increasing ozone or, in one case, stayed roughly constant. The  
concentration of the meteoric-sulfuric particles is fairly consistent between hemispheres and deployments. In contrast, the  
concentration of organic-sulfate particles was larger in the Northern Hemisphere than in the Southern Hemisphere and varied  
considerably between deployments. It is worth noting that the highest concentration of tropospheric particles in each  
hemisphere, ATom3 for the Southern Hemisphere and ATom4 for the Northern Hemisphere, were both local springtime.

240

Figure 6 shows the ratio of the C<sup>+</sup> peak, an indicator of organic content, to two peaks indicative of sulfate or sulfuric acid. The  
top axis gives an approximate mass fraction of organics adapted from calibrations described by Froyd et al. (2019). The vertical  
axis of ozone serves as a measure of distance into the stratosphere. The organic content is separated by particle type, something  
not possible with bulk analysis. That the stratospheric and tropospheric particle compositions remain distinct implies that there  
is very limited redistribution of semi-volatile organics between particles. Like most upper tropospheric particles, the organic-  
sulfate particles are internally mixed with ~~contain~~ on average about 40 to 80 percent organic material by mass. There is little  
variation with ozone, indicating a long lifetime for the organic material as well as little uptake of sulfuric acid. The latter is  
consistent with most of the sulfuric acid coming from carbonyl sulfide above 20 km rather than SO<sub>2</sub> near the tropopause  
(Kremser et al., 2016; Rollins et al., 2017). That there are two types of particles with different compositions present in the  
same air also means that any semi-volatile gas-phase organics and ammonia cannot be in equilibrium with both types of  
particles.

There is some limited uptake of organics onto the stratospheric particles, although the maximum organic concentration is still  
much less than for tropospheric particles. Meteoric-sulfuric particles definitely formed in the stratosphere, so any significant  
organic content indicates net uptake of organics. Their organic content grows as the particles descend to the lowermost  
stratosphere and the upper troposphere. The meteoric-sulfuric particles contain much less than 1% organic mass at altitudes  
with ozone greater than 500 ppbv and as much as 2 to 4% near the tropopause in the Northern Hemisphere. Such limited  
formation of secondary organic mass in the lowermost stratosphere is consistent with previous PALMS measurements  
(Murphy et al., 2007).

260

A new finding from ATom is that there is a very distinct and consistent difference between the hemispheres in the small amount  
of organic content that does form in the meteoric-sulfuric acid particles. Since the particles start from similar formation  
processes much higher in the stratosphere, we conclude that there is more condensable or reactive organic vapor in the Northern  
Hemisphere lower stratosphere. This could be either gas phase species mixed from the troposphere or semi-volatile organics  
transferring from organic-sulfate particles.

Adding a few percent mass to the meteoric-sulfuric particles represents a very small amount of organic vapor. Without knowing  
uptake coefficients the amount of vapor cannot be determined uniquely, but a representative calculation is that one or two pptv

of an organic gas-phase species with molecular weight of about 100 daltons that reacts with sulfuric acid on every collision would add few percent mass to a 450 nm particle in a few months. The same order of magnitude can be obtained by noting from Figure 5 that at 200 ppbv ozone there is about 100 ng standard m<sup>-3</sup> of meteoric sulfuric particles. 1% by mass of these particles corresponds to about a part per trillion by mass of air. We conclude that an order of magnitude for highly condensable organic vapor in the lowermost stratosphere is a few parts per trillion in the Northern Hemisphere and less in the Southern Hemisphere. A less reactive or condensable organic molecule could be present at a correspondingly higher concentration.

## 5 Radiative and chemical implications

The different sizes of the sulfuric acid and organic-sulfate particles lead to substantial differences in their radiative and chemical effects. Important properties are the amount of infrared heating, the amount of light scattered, implications for photolysis, the surface area available for heterogeneous chemistry, and the mass sedimentation rate.

A key part of the radiative implications is the efficiency of light scattering as a function of particle size. Figure 7 shows the mass scattering efficiency as a function of particle size averaged over the solar spectrum and a sunlit Earth. Calculations are for a real refractive index of 1.45 and minimal absorption. Atmospheric extinction is determined by the solid total scattering curve. Much of the light scattered by particles continues downward to the Earth and so does not directly affect climate. Separating out the light scattered to outer space (dashed curve) gives a maximum that is slightly broader and shifted to smaller sizes than light extinction. Over much of the size range of particles that scatter light efficiently, only about 1/5 of the light that is scattered goes to outer space; the remainder becomes diffuse light. This is a reason for the large increases in diffuse light (with decreases in direct sunlight) after volcanic eruptions (Murphy, 2009).

The top panel of Figure 8 replots the chemically resolved size distribution for ATom2 in the Northern Hemisphere. This was chosen as an example because it is similar to several other locations and seasons such as ATom1 Northern Hemisphere and ATom3 Southern Hemisphere. There was more sampling time in the ATom2 Northern Hemisphere stratosphere so the particle statistics are better, with about 10,000 particles. In Figure 8 the two sulfuric acid particle categories have been combined. Of the particles larger than 0.1  $\mu\text{m}$  diameter, about 39% of the volume was organic-sulfate particles from the troposphere and 61% sulfuric acid particles from the stratosphere (including both those with and without meteoric metals). The percentage contribution to each parameter in Figure 8 by stratospheric aerosol will be somewhat larger at ambient conditions because sulfuric acid has some water uptake even at <10% relative humidity. Ambient sulfuric acid particles may have roughly 5 to 15% larger diameters than measured in the warm aircraft cabin.

### 5.1 Infrared heating

An important property for stratospheric particles is their absorption and emission of infrared radiation. Infrared properties are more important for stratospheric aerosol than very low-altitude aerosol because the latter are close to the surface temperature and so absorb and emit similar amounts of energy.

Infrared absorption by stratospheric particles is important for two reasons. First, it heats the stratosphere around the particles. Changes in circulation due to infrared heating were responsible for significant changes in ozone after the Pinatubo eruption (Labitzke and McCormick, 1992; Pitari and Rizi, 1993). The heating-induced changes in ozone were as large or larger than those due to heterogeneous chemistry. There are additional feedbacks on the circulation after changes in infrared heating due to changes in ozone and water vapor (Visioni et al., 2017). Infrared heating is largest in the lower stratosphere where the temperature contrast with the surface is greatest (Lacis, 2015).

Second, infrared absorption by stratospheric particles offsets some or even all of the shortwave cooling of the Earth. For sulfuric acid particles similar to those after the Pinatubo eruption, longwave heating offset roughly 25% of the shortwave cooling (Hansen et al., 2005). This increases to about 50% for large injection rates because larger particles ( $>0.6\text{ }\mu\text{m}$ ) become increasingly less efficient at scattering sunlight to outer space (Figure 67) compared to their volume (Niemeier and Timmreck, 2015). Large particles can even cause net warming: alumina with the size distribution from rocket emissions was calculated to cause net warming (Ross and Sheaffer, 2014).

For wavelengths much larger than the particles, absorption and emission are approximately proportional to total particle volume (and the material) and do not depend on particle size (van de Hulst, 1981). Therefore, net thermal infrared heating is insensitive to the size of particles and will approximately follow the volume distribution in the top panel of Figure 78. For ATom2, most of the volume distribution was composed of particles of stratospheric origin, but that was not always the case (Figure 1). The infrared effects of the tropospheric particles found in the lower stratosphere are hard to assess. Longwave radiative heating depends not only on the strength of the absorption bands but also on their overlap with atmospheric windows in the infrared spectrum. The presence of a significant fraction of organic material has unknown implications for infrared heating. Like sulfate, infrared heating will be proportional to the volume of organic aerosol, but the absolute amount of heating depends on how aerosol absorption features correlate with gas-phase absorption or window regions in the infrared spectrum. We are not aware of any infrared spectra of organic material in aerosols in the stratosphere or upper troposphere.

## 5.2 Scattering of visible light

The middle panel in Figure 7-8 shows the size distribution weighted by the amount of sunlight scattered to outer space, which is relevant for shortwave climate effects. Weighting the size distribution by the light extinction would shift the peak just slightly further to larger particles. Because of their size, the sulfuric acid particles contribute a greater fraction of the light scattering

than their mass fraction. In fact, comparing Figures 1 and 67, the most abundant size of the sulfuric acid particles in the lower stratosphere was close to the maximum in light scattering to outer space per unit mass. The ATom mission was during a time with small volcanic influence. In contrast, particles shortly after the Pinatubo eruption had volume mean diameters greater than 0.7  $\mu\text{m}$  (Wilson et al., 2008; Figure S2S3), large enough that their mass scattering efficiency decreased. [Figure S4 shows the relative extinction due to various particle types for cases other than the ATom2 Northern Hemisphere used as an example in Figure 8.](#)

### 5.3 Heterogeneous chemistry

Particle size can affect heterogeneous chemistry. Reactions with sulfuric acid particles that are important to stratospheric chemistry span the range from reactions that occur in the interior of liquid particles and hence are proportional to volume to reactions that occur on the surface and hence are proportional to surface area (Hanson et al., 1994). Heterogeneous chemistry can be especially important within or at the edge of the polar vortex (Solomon et al., 2015; Stone et al., 2017).

The bottom panel of Figure 8 shows the size distribution weighted by surface area rather than volume. For these conditions surface reactions are closely proportional to surface area; gas-phase diffusion is a minor correction. For the case shown in Figure 8, the organic-sulfate particles from the troposphere are about half of the surface area in the lowermost stratosphere. This is significant because whereas stratospheric heterogeneous chemistry on sulfuric acid has been extensively studied, little is known about the same reactions on organic-sulfate particles. [Figure 8 shows ATom2 Northern Hemisphere as an example. Surface area fractions for other cases are shown in Figure S4.](#)

The organic-sulfate particles differ from the sulfuric acid particles in important ways. Most obviously, they contain a high proportion of organics that may participate in new chemistry with halogen radicals. Iodine in particular may react with organic aerosols (Murphy and Thomson, 2000; Pechtl et al., 2007). Although not fully neutralized, the organic-sulfate particles are not nearly as acidic as the relatively pure sulfuric acid particles. This can be determined from acid cluster peaks in the PALMS mass spectra. Some chlorine activation reactions that lead to ozone destruction are acid-catalyzed (Burley and Johnston, 1992) and therefore may be slower on partially neutralized particles. The organic-sulfate particles also contain less water – sulfuric acid is extremely hygroscopic compared to other species at the low relative humidities in the stratosphere. The availability of condensed water for heterogeneous reactions could be further reduced if the organic-sulfate particles are glassy at the low temperatures and humidity in the stratosphere (Krieger et al., 2012; Price et al., 2014).

### 5.4 Sedimentation

370 Sedimentation is a key process in the stratospheric aerosol budget (Wilson et al., 2008). It is more important than in the  
stratosphere than it is near ground level partly because particles fall faster at lower air density. A bigger reason sedimentation  
is important in the stratosphere is the relevant time scale: a fall speed of a kilometer per month would be unimportant in the  
lower troposphere but can control the residence time of a particle in the stratosphere. Particles larger than about 1  $\mu\text{m}$  diameter  
have sedimentation rates greater than 10  $\text{km yr}^{-1}$  in the lower stratosphere. ~~On Figure 7, the The~~ sedimentation flux as a function  
375 of size, ~~shown in Figure S5, is similar to the light scattering panel in Figure 8 would be slightly more skewed to large diameters  
than the light scattering panel).~~ For the case shown in Figure 7 For the ATom2 example used in Figure 8, the sulfuric acid  
particles have roughly twice the volume of the organic-sulfate particles. Their source strength must be somewhere between  
about twice as large as the tropospheric particles (if loss is controlled by bulk air motion) and three times as large (if loss is  
controlled by sedimentation).

380

## 5.5 Ultraviolet scattering and absorption

Absorption and scattering of ultraviolet light are distinct from that of visible light because of the impact on photolysis rates.  
Light scattering by stratospheric aerosol changes the path length of light in the stratosphere, which in turn changes photolysis  
385 rates (Huang and Massie, 1997; Pitari et al., 2015). The calculations are complex because Rayleigh scattering in the ultraviolet  
leads to strong effects from multiple scattering (Bian and Prather, 2002), especially at twilight or if a scattering aerosol layer  
is located above gas-phase absorption (Davies et al., 1993; Anderson et al., 1995). In addition, the long path lengths magnify  
the importance of any absorption of ultraviolet light by aerosols.

390 The scattering of ultraviolet light peaks at smaller particle sizes than for sunlight. This means that the smaller tropospheric-  
organic particles contribute substantially to the scattering of ultraviolet light. The relative contributions to scattering of light  
at  $< 240 \text{ nm}$  are ~~fairly similar~~ shown in Figure S5 and are similar to the surface area panel in Figure 8 except that sizes smaller  
about 80 nm and larger about 600 nm contribute less to UV scattering than they do to surface area.

395 One important wavelength band is 200 to 242 nm, where photolysis of  $\text{O}_2$  is responsible for formation of ozone and photolysis  
of  $\text{N}_2\text{O}$  produces odd nitrogen ( $\text{NO}_y$ ). (Brasseur and Solomon, 1986). For purely scattering particles, changes in photolysis in  
this wavelength range are reduced by large cancellations in direct and diffuse light (Michelangeli et al., 1989). Light scattering  
by the El Chichon volcanic cloud was estimated to reduce  $\text{O}_2$  photolysis by about 10% (Michelangeli et al., 1989). The overall  
effect of scattering seems to be to reduce ozone formation (Pitari et al., 2015).

400

Unlike pure scattering particles, absorbing particles would not have a similar partial cancellation between direct and diffuse  
sunlight. Huang and Massie (1997) examined the effect of substituting ash, with visible and UV absorption, for non-absorbing  
sulfuric acid in a simple model of photolysis after a volcanic eruption. There are competing effects on the ozone column

because both  $J_{O_2}$  and  $J_{O_3}$  are reduced by UV absorption, with one reducing and the other increasing ozone (Pitari and Rizi, 1993). The individual effects were several percent of the ozone column with the net impact difficult to assess because their simple model did not include  $NO_x$  or halogen chemistry. The imaginary refractive index of the organic-sulfate particles at wavelengths below 242 nm is not known but could easily be large enough to lead to significant absorption compared to scattering.

## 6 Relevance to volcanic or intentional aerosol injection

Sulfur dioxide is periodically injected into the stratosphere by major volcanic eruptions, leading to an increase in sulfuric acid aerosol. One such example is visible in the ATom1 Southern Hemisphere panel in Figure 1. There have also been numerous proposals for, and studies of, injecting material into the stratosphere for the purpose of solar radiation management (National Research Council, 2015). Regardless of the desirability of such actions, some of the calculations presented above on the optical properties and potential for heterogeneous chemistry have implications for added material whether it be from volcanoes or deliberate intervention. Even if various types of particles have similar optical properties, the chemically resolved size distributions are important for attributing radiative effects to different sources. Without information about particle composition we could not say, for example, what percentage of extinction is due to particles formed in the stratosphere and what percentage is due to particles formed in the troposphere.

### 6.1 Complex controls on particle size

The previous sections demonstrate that particle size is important for many of the properties relevant to climate and chemical effects. In the real world, one cannot instantly fill a box with monodisperse particles of a chosen size, the way one might in a model. It is important to understand what controls the size distribution of particles in the stratosphere and how it will~~will it~~ change with additional aerosol or precursors such as sulfur dioxide. The mean diameter in the stratosphere is not a single number but varies with the aerosol loading, altitude, and latitude (English et al., 2012).

The size of particles in the unperturbed or perturbed stratosphere can be understood in two complementary ways. The first way, a top-down approach, says that for a given mass of stratospheric aerosol, the more particles there are, the smaller they must be. The second way, a bottom-up approach, considers how the size of each particle is set by a balance of growth and removal processes in the stratosphere.

For the top-down approach, one must consider at least three sources of particle number in the stratosphere. Particles come down from the mesosphere, up from the troposphere, and new particles can form in the stratosphere (Murphy et al., 2014). The meteoric source of particles to the stratosphere is mostly “smoke” consisting of material that evaporated from ablating



meteoroids and condensed into new particles high in the atmosphere. Much of this material descends near the winter poles (Bardeen et al., 2008). Second, tropospheric particles provide an important source of stratospheric particles below 20 km altitude (Yu et al., 2016). The fate of tropospheric particles entering the stratosphere is poorly represented in most models. New particle formation is also important for the stratosphere. The pure sulfuric acid category in Figure 1 is probably from growth of particles formed in (or at the edge of) the stratosphere. One formation region is near the tropical tropopause with upward transport into the stratosphere (Brock et al., 1995; English et al., 2011). There is probably also formation of new sulfuric acid particles higher in the stratosphere over the winter poles (Wilson et al., 1989), although this must be distinguished from meteoric smoke descending in the same regions (Curtius et al., 2005).

Adding sulfuric acid or its precursors will have complex effects on new particle formation, with more vapor to condense but also more surface area sink. In contrast, injected solid particles would provide a surface sink for background sulfuric acid from oxidation of carbonyl sulfide, likely reducing new particle formation. This implies that injected solid particles would probably change the size of the natural sulfuric acid particles in the stratosphere. Sufficiently small injected solid particles might reach high altitudes where existing sulfuric acid particles have evaporated (Weisenstein et al., 2015; Jones et al., 2016). There could be unknown effects if they were later entrained in descending air in the winter polar regions.

The bottom-up approach considers how the size of stratospheric particles is determined by a balance of growth and removal processes. Particles grow by coagulation and by condensation of sulfuric acid and other species. Coagulation in the unperturbed stratosphere is slow except for special situations such as shortly after new particle formation (Brock et al., 1995; Hamill et al., 1997). Coagulation increases non-linearly with aerosol concentration so it becomes more significant after volcanic eruptions (Pinto et al., 1989) or large injection scenarios (Weisenstein et al., 2015). In these cases, coagulation helps drive the extra mass primarily to larger particles rather than more numerous particles (Heckendorn et al., 2009; Niemeier and Timmreck, 2015). Both sedimentation and downward motion are important removal processes (Wilson et al., 2008).

One implication of having multiple sources of particles in the stratosphere is that there is no single response to injected material. It is only in the last few years that stratospheric models have incorporated multiple sources of particles along with detailed microphysics (Pitari et al. 2014; Yu et al., 2016; Mills et al., 2017). There is still considerable uncertainty in quantitatively understanding the size of particles in the current stratosphere, let alone after a perturbation. Figure 2 demonstrates that a detailed microphysics model of the stratosphere did not grow the sulfuric acid particles to large enough sizes.

## 6.2 Impacts on the stratosphere as a function of size

The preceding section suggests that the ultimate size of particles is not something that can simply be chosen when injecting aerosol particles or their precursors; instead it is also strongly influenced by subsequent particle formation and growth

processes in the stratosphere. Even solid particles may coagulate or grow by condensation of sulfuric acid and not retain their original size. Despite this, it is of interest to examine the impact of particle size on the relative magnitude of effects. Figure 9 shows an estimate of net cooling compared to particle volume, surface area, and sedimentation rate.

The shortwave effect is modified to include longwave effects by adapting a calculation in Lacis et al. (1992), who showed that heating exceeded shortwave cooling for several different types of particles when their diameters were more than about 4  $\mu\text{m}$ . Since infrared absorption is almost independent of size, an approximate net cooling was estimated by subtracting a constant from the calculated shortwave scattering to space (Figure 6) such that the net was zero at 4  $\mu\text{m}$  diameter. Figure 7 shows the infrared effect on net cooling as a function of size. The vertical scale of the net cooling curve is consistent with the scattering curves. For example, at peak scattering efficiency near 0.5  $\mu\text{m}$ , infrared effects reduce the cooling of the Earth by stratospheric aerosol by 10 to 15%. The relative infrared effect is larger for either larger or smaller particles. Subtracting this amount This calculation is only approximate but the patterns in Figure 9 are not sensitive to the exact crossing point amount subtracted from cooling by pure scattering. The relative sedimentation curve in Figure 9 also varies slightly with altitude but again the patterns are similar.

The size that has the least infrared absorption for a given amount of shortwave climate impact is about 0.5  $\mu\text{m}$  diameter (filled circles). This is true regardless of details of the infrared spectrum of the particles. Larger or smaller particles will be less effective at cooling the Earth and will cause more stratospheric circulation changes for a given amount of cooling. Sufficiently large ( $> 4 \mu\text{m}$ ) or small ( $< 0.1 \mu\text{m}$ ) particles cause net heating of the Earth (Lacis et al., 1992). The crossing point between net cooling and net heating will depend on particle composition. Even for the optimal size, the infrared heating due to deliberate injections of sulfuric acid or its precursors into the stratosphere would cause significant changes in circulation (Aquila et al., 2014).

The potential for increased heterogeneous chemistry would be reduced by using larger particles with less surface area. Particles with a diameter of about 1  $\mu\text{m}$  have the largest cooling effect for a given surface area (open circles). These larger particles, however, have high sedimentation rates (downward triangles) and limited stratospheric lifetimes. For a given climate impact, the mass flux due to sedimentation is minimized by particles with a diameter of about 300 nm. The least diffuse light is created by the smallest particles (Figure 7). The exact sizes of the various maxima depend on the refractive index and density. However, the patterns of which variables are maximized at larger or smaller sizes will be similar.

## 7 Discussion

The title of Solomon et al. (2011) includes the phrase “the persistently variable ‘background’ stratospheric aerosol”. The ATom data presented here add new meaning to that phrase. Out of eight samplings of the lowermost stratosphere, three exhibited

much higher aerosol concentrations for three different reasons: a volcanic eruption, biomass burning aerosol, and transport of dust and other near-surface particles. None of these events were targeted during the flights. Such variations in the stratospheric aerosol layer are important for both heterogeneous chemistry (Solomon et al., 1996) and climate (Solomon et al., 2011).

510 There are important differences in the aerosol in the lower stratosphere between the Northern and Southern Hemispheres. A smaller amount of tropospheric aerosol in the Southern Hemisphere stratosphere indicates that the tropospheric particles are mixing into the lower stratosphere within each hemisphere rather than entering in rising air in the tropics and splitting into the two hemispheres. Sulfuric acid particles in the Southern Hemisphere also acquire less organic content. This suggests that there are lower concentrations of gas-phase organics in the Southern Hemisphere. One of several possible formation routes is that

515 small organic compounds such as acetone and formaldehyde can react with concentrated sulfuric acid to form polymers that stay in the aerosol (Iraci and Tolbert, 1997; Williams et al., 2010). Other routes would be if low-volatility organic molecules were formed in the gas phase or evaporated from the tropospheric particles and recondensed on the sulfuric acid particles. Even in the Northern Hemisphere, only low part-per-trillion range concentrations of gas-phase organics are required to explain the very small amounts of organics taken up by the sulfuric acid particles.

520 The data here add support to the concept of Yu et al. (2016) that tropospheric particles comprise a significant fraction of the aerosol in the lowermost stratosphere. Such tropospheric particles offer a route for anthropogenic influence on the stratosphere. The Yu et al. model also correctly predicts that tropospheric particles are smaller than sulfuric acid particles formed in the stratosphere (Figure 2).

525 The data here have several implications for satellite retrievals. First, the reason lower stratospheric size distributions are often broad is that they are really two or more overlapping distributions. As the tropospheric/stratospheric mix shifts the mean diameter will shift. Second, unlike sulfuric acid, the smaller, organic-rich particles coming up from the troposphere may be effloresced and/or glassy. That could explain a very small amount of depolarization. Finally, when only the stratospheric mode

530 is dominant the size distribution can be very narrow.

Absorption of ultraviolet light means that impurities should be considered when assessing deliberately added materials. For example, absorption appropriate for optical-quality sapphire should probably not be used when evaluating proposals to add industrial quantities of alumina to the stratosphere. Even part-per-million impurities in alumina increase absorption in the

535 ultraviolet (Innocenzi et al., 1990). Compared to many materials, sulfuric acid has extremely low absorption in the ultraviolet (Noziere and Esteve, 2005; Dykema et al., 2016).

The broad distribution of particle sizes in the unperturbed stratosphere is the superposition of several narrower distributions. Single particle types, particularly meteoric-sulfuric acid particles, can have narrow size distributions (lognormal standard deviation ~1.4).

Multiple formation mechanisms for stratospheric particles imply that the size of particles after a volcanic or intentional injection may be difficult to predict. Yet an accurate prediction of size is important: The diameter must be known to perhaps 25% to accurately estimate tradeoffs between climate impact and side effects (Figure 9). A state-of-the-art microphysical bin model underestimates the size of stratospheric sulfuric acid particles, indicating that we do not fully understand what controls the size of particles in the stratosphere. The size difference has significant impacts on properties: the modeled particles have about 65% of the climate impact per unit mass as ~~the observations~~ calculated from observed sizes of stratospheric sulfuric acid particles, 160% of the surface area, and sediment about 60% as fast.

There is no single diameter that produces the largest shortwave climate impact with the fewest side effects (Figure 89). To the extent that one could control the size of particles after an intentional injection, any chosen size involves tradeoffs. Particles smaller than about 0.6  $\mu\text{m}$  diameter have more surface area for possible heterogeneous chemistry. Particles larger than about 0.4  $\mu\text{m}$  require more injected mass and produce more diffuse light. For a given amount of scattered sunlight, either sufficiently large or small particles have more infrared absorption and hence more impacts on stratospheric circulation. Most of the mass of particles after the Mt. Pinatubo eruption was larger than 0.6  $\mu\text{m}$  diameter (Brock et al., 1993; Wilson et al. 2008), a size range with relatively little surface area compared to their climate impact. The heterogeneous chemistry observed after Mt. Pinatubo may therefore underestimate what might happen with intentionally added material.

#### Data availability and author contributions

Data are available at <https://daac.ornl.gov/ATOM/campaign/>, <https://espo.nasa.gov/atom/content/ATom>, and <https://www.esrl.noaa.gov/csl/projects/atom/data/MurphyACP2020/>. DMM wrote the manuscript with assistance from all other authors. DMM, KDF, CAB, AK, GPS, and CJW made aerosol measurements and analysed the data. IB and CRT made ozone measurements and analysed the data. PY performed model calculations. Authors declare no competing interests.

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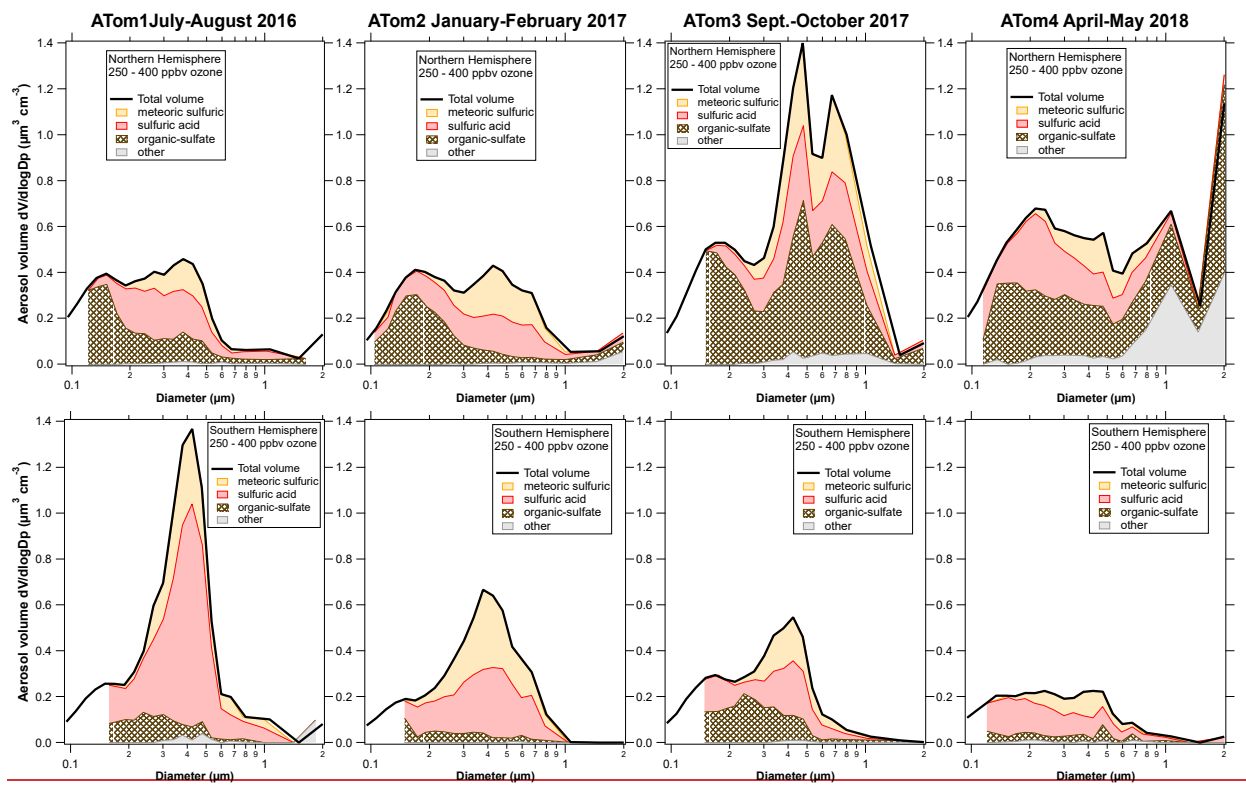
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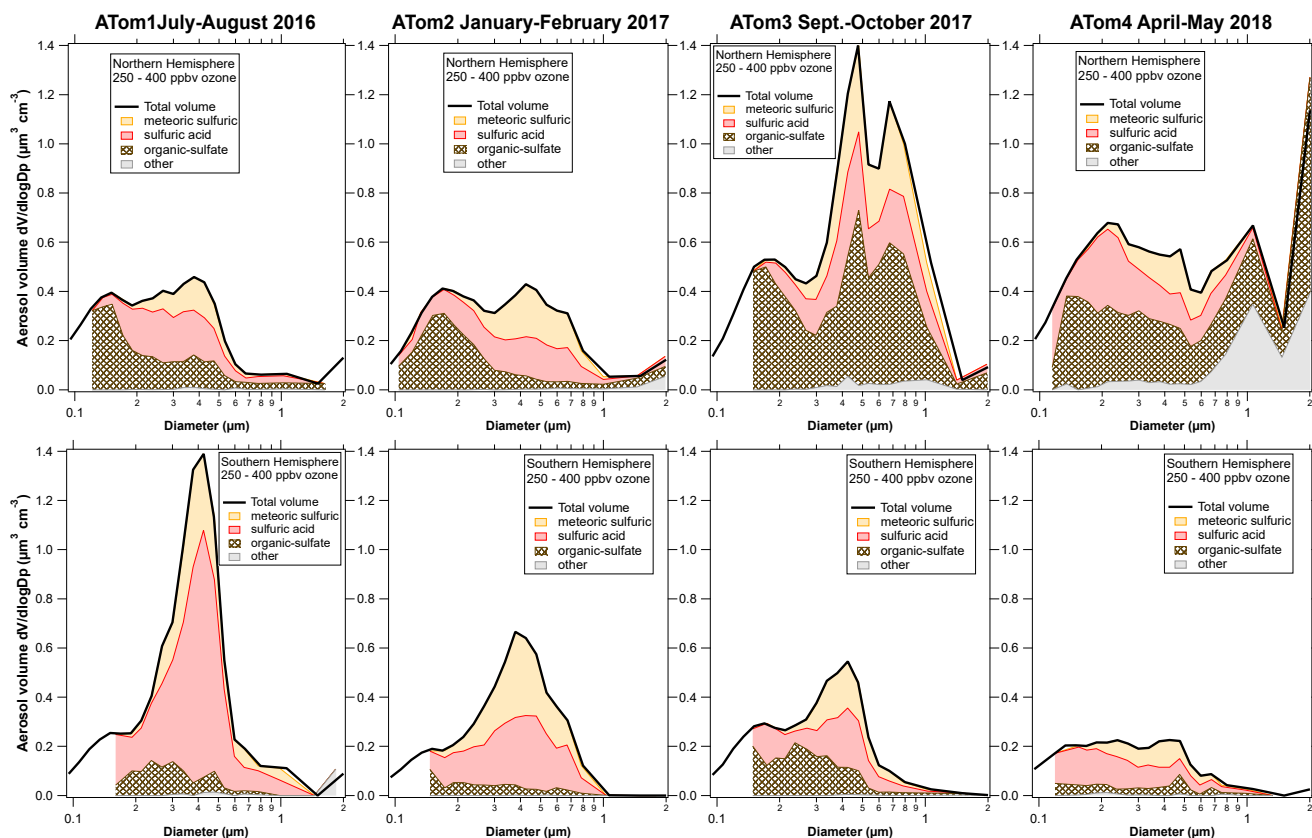
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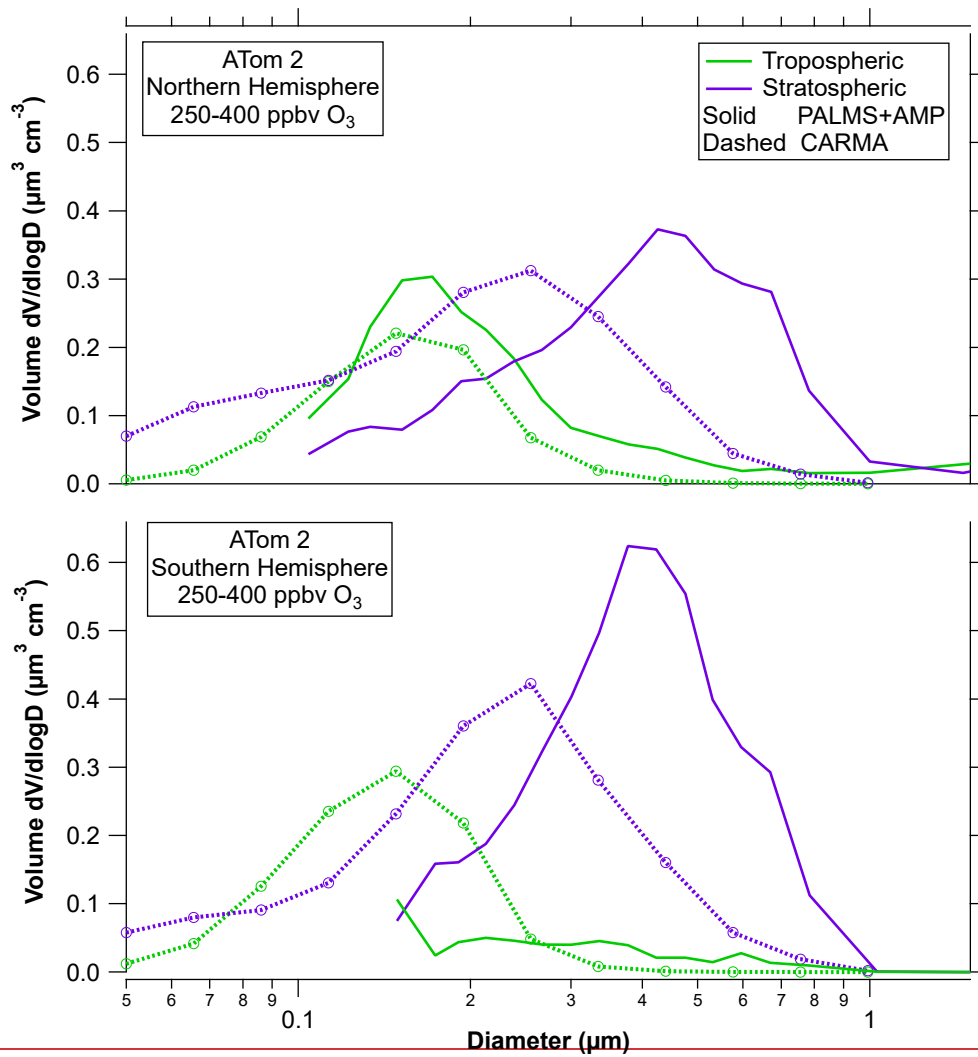




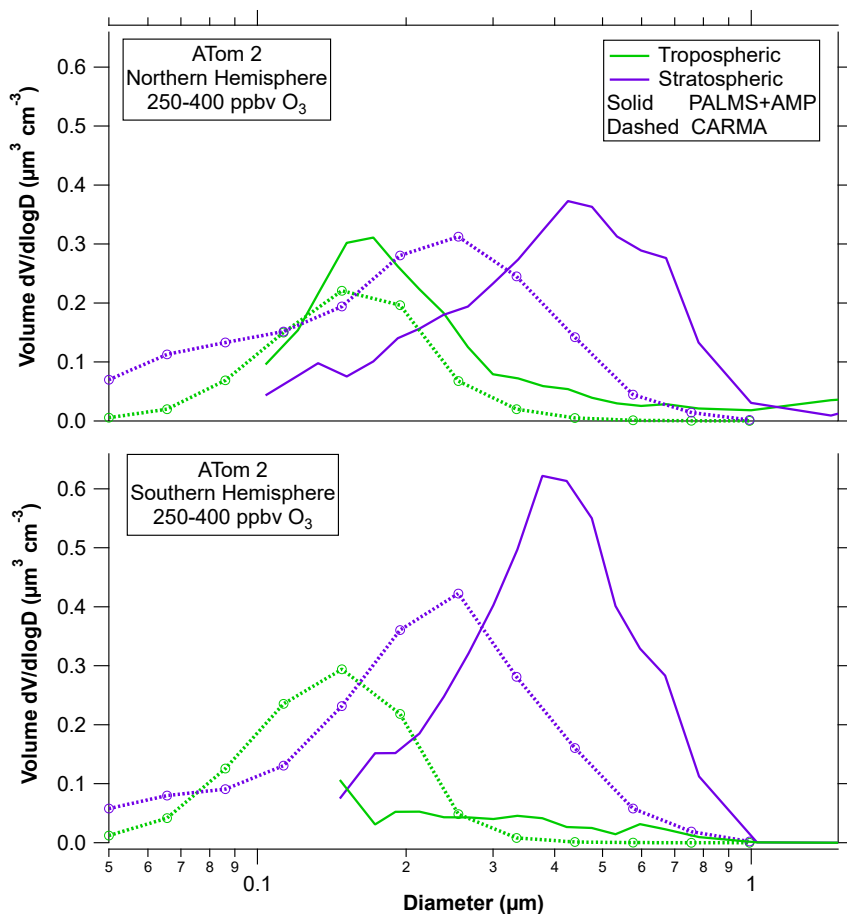


**Figure 1. Composition resolved size distributions from the lower stratosphere during the Atmospheric Tomography (ATom) mission. The size distributions from optical particle counter measurements are apportioned at each size into classes of particles based on the PALMS single particle composition data. The smallest few size bins sometimes had too few particles with PALMS mass spectra to apportion the composition. Sizes in all figures are geometric diameter at low relative humidity and concentrations are per standard  $\text{cm}^3$ . Fine structure in some total volume distributions at  $0.6 \mu\text{m}$  and larger may be artifacts due to Mie resonances in the optical particle counter; minima near  $0.2$  or  $0.3 \mu\text{m}$  are robust.**



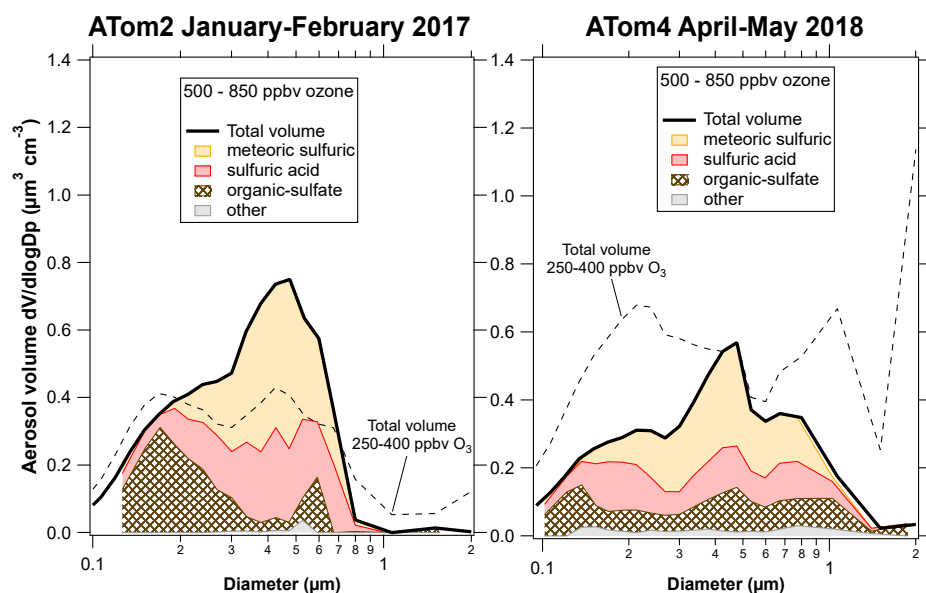


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**Figure 2. A comparison of the CARMA bin microphysics aerosol model in the CESM climate model with observed size distributions during ATom2. Total particle volumes are proportional to areas under each curve.**

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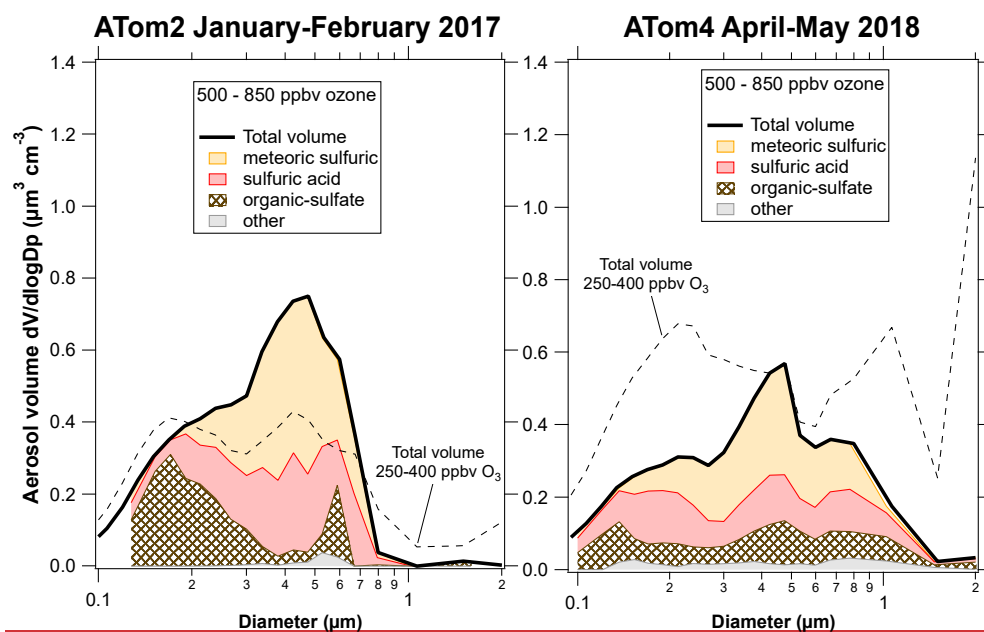
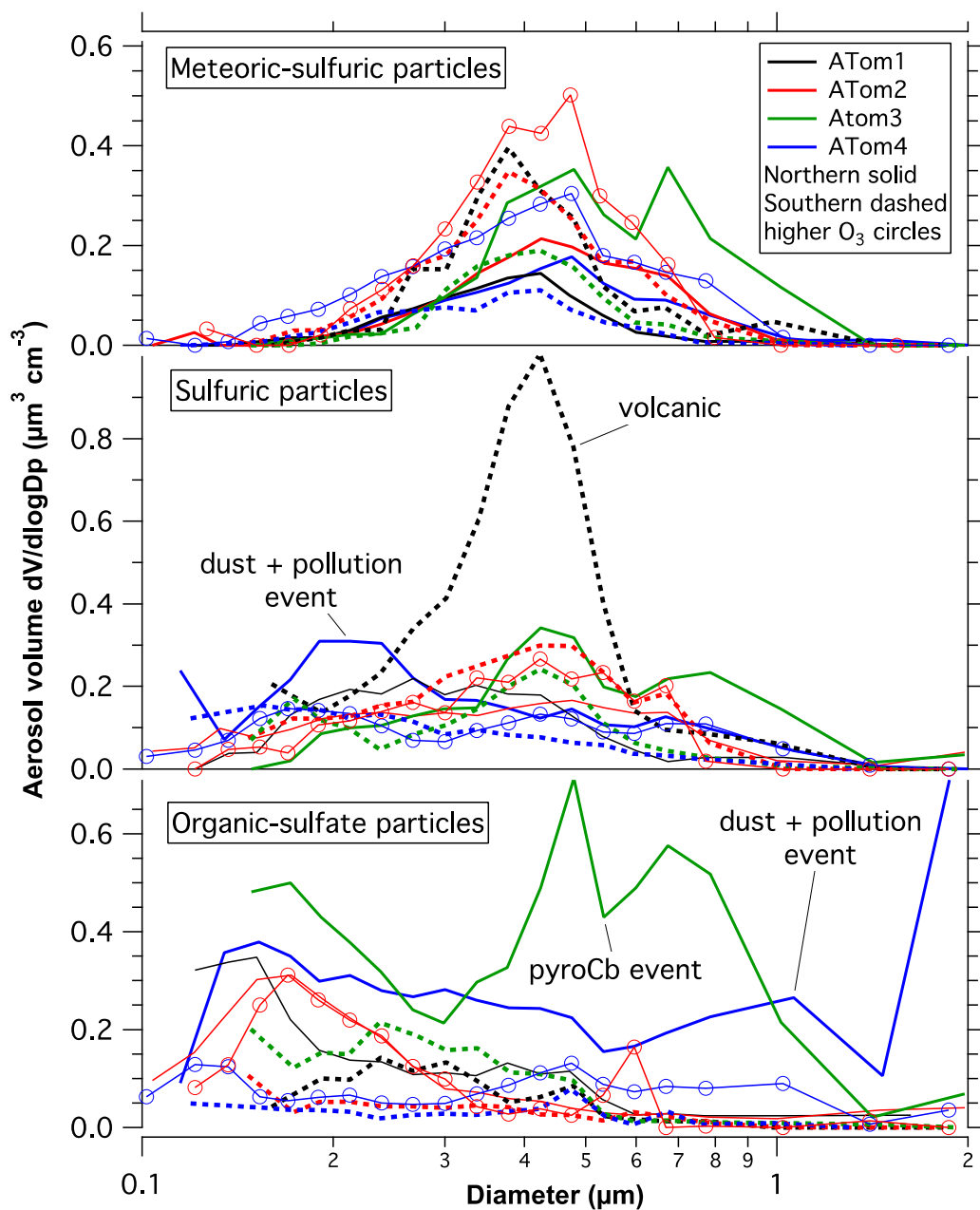


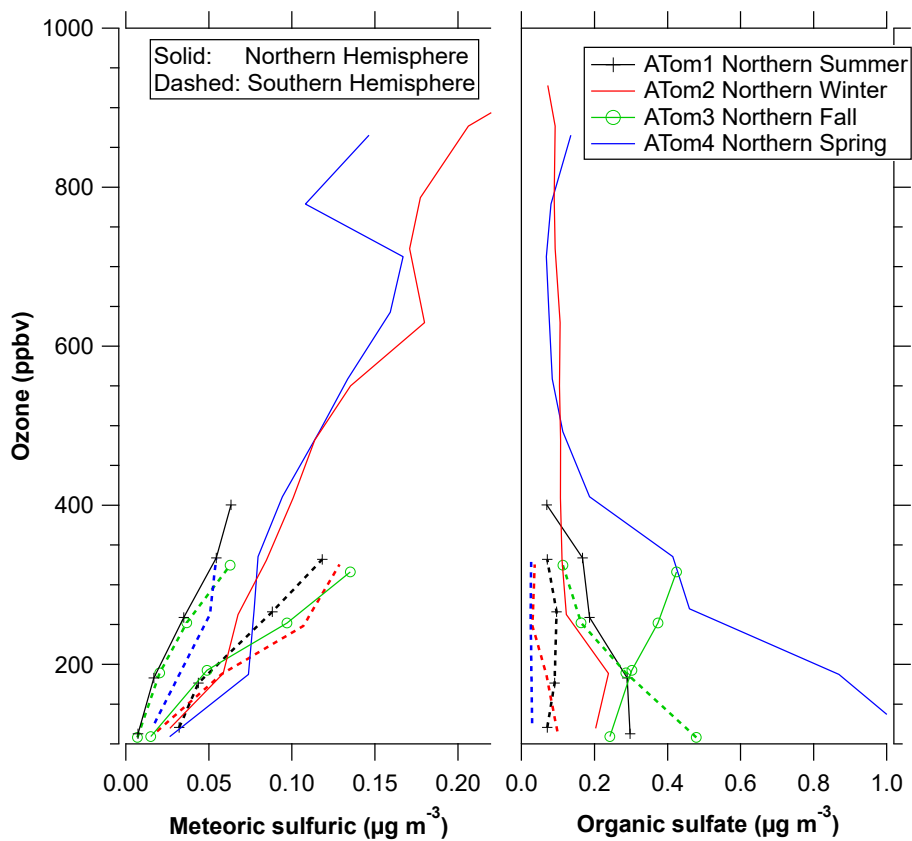
Figure 3. Composition-resolved size distributions for the times when data are available for ozone greater than 500 ppbv, all in the Northern Hemisphere. The dashed curves are to facilitate comparison with Figure 1.



**Figure 4.** Size distributions of several types of particles in the lower stratosphere. Northern Hemisphere data are solid and Southern Hemisphere data are dashed. Data are from 250 to 400 ppbv ozone except the curves with circles are data at more than 500 ppbv ozone. The three major events are noted where they may have influenced the size distributions. Lognormal fits are given in Table 1.







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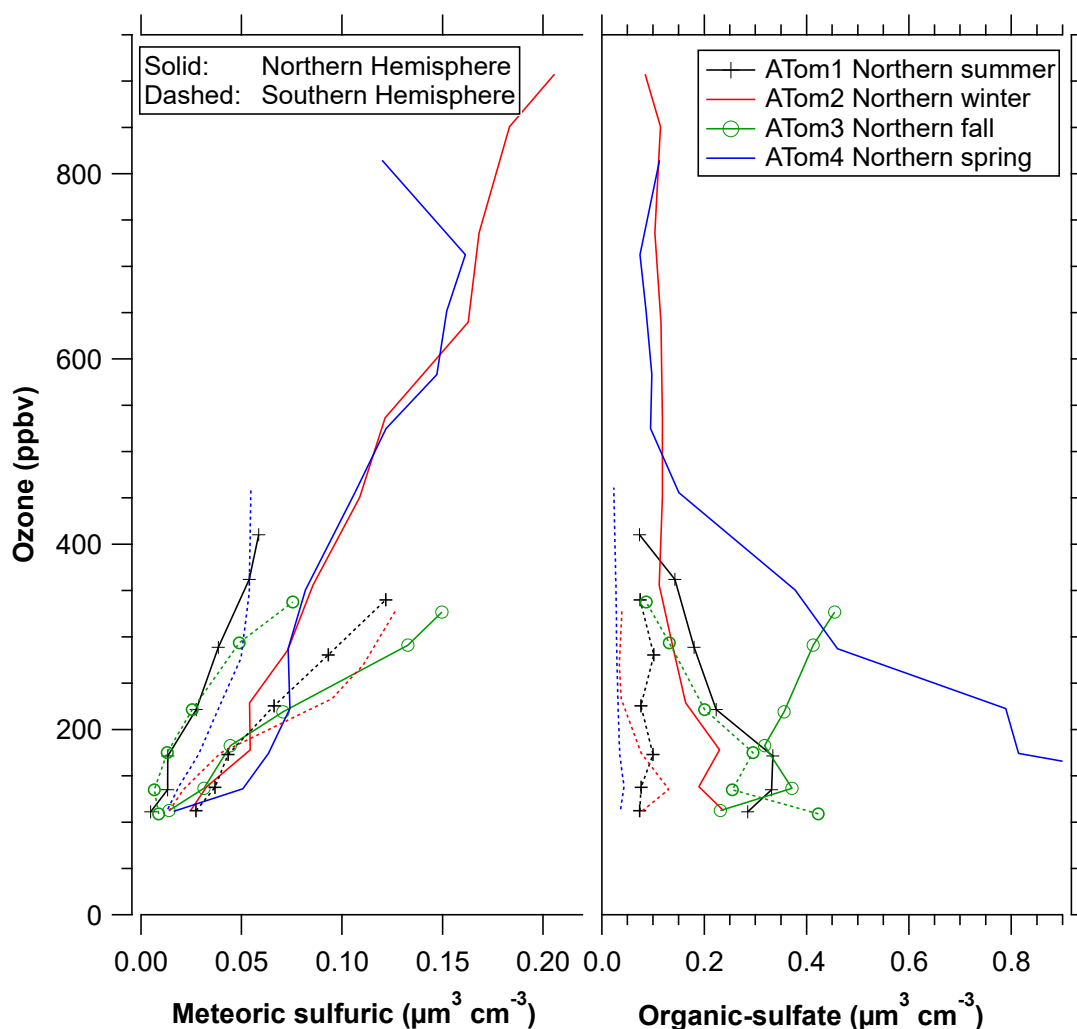
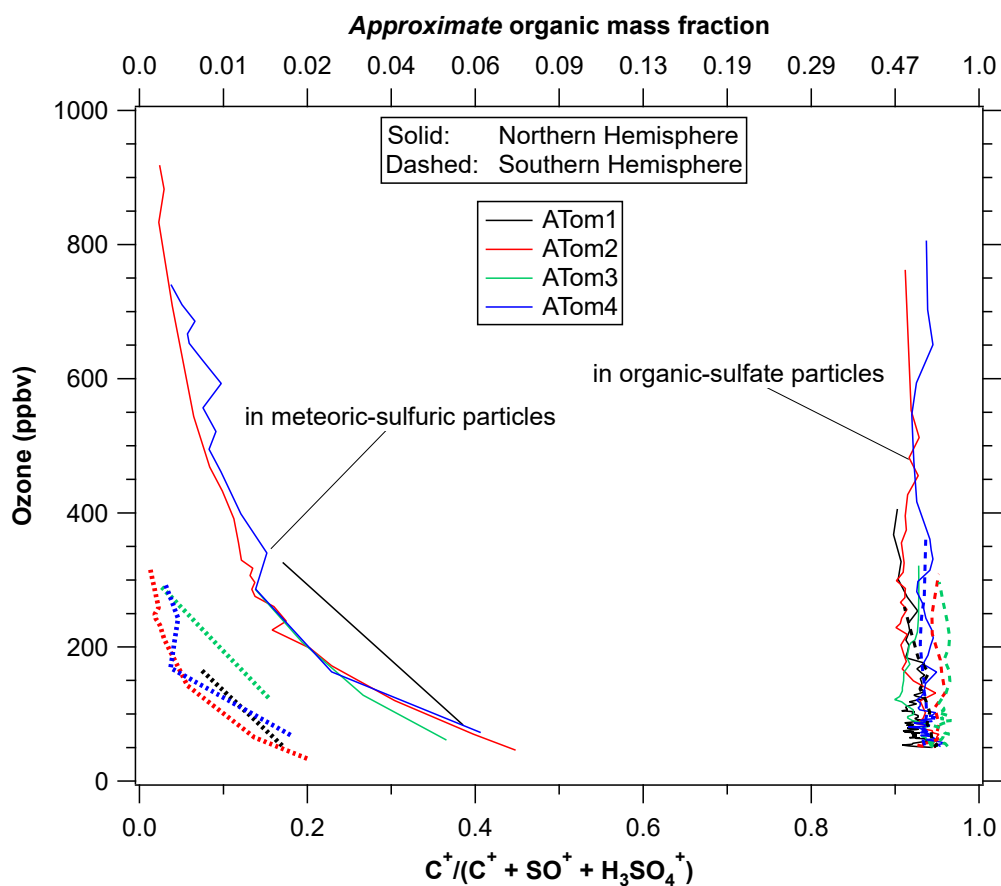


Figure 45. Profiles of the mass-volume concentration of two types of particles in the lower stratosphere, using ozone as a surrogate for a vertical coordinate. This does not include particles smaller than 0.1  $\mu\text{m}$ . Sulfuric acid particles with meteoric metals have a high-altitude source whereas organic-sulfate particles, as selected by their mass spectra, originate in the troposphere. The organic-sulfate particles are more abundant in the Northern Hemisphere than they are in the Southern Hemisphere.





905 **Figure 56.** The organic content of two types of stratospheric particles using ozone as a surrogate for altitude. Shown  
 are the relative areas of the  $C^+$  peak in the mass spectra relative to two peaks indicative of sulfate. The top axis is an  
 approximate mass fraction. The organic-sulfate particles are about 40 to 70% organic by mass, consistent with a  
 tropospheric source. The sulfuric acid particles have near-zero organic content well into the stratosphere, increasing  
 up to a few percent by mass near the tropopause. There was a strong difference in the organic content of the sulfuric  
 910 acid particles between the hemispheres. Each point is the average of about 200 mass spectra.

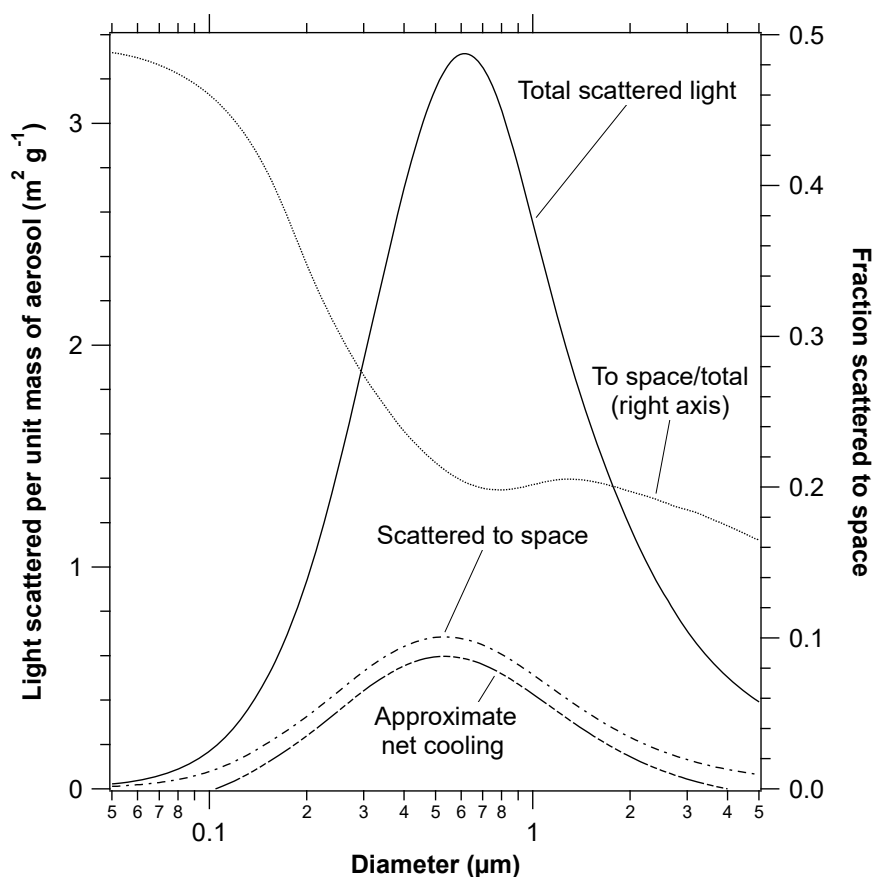
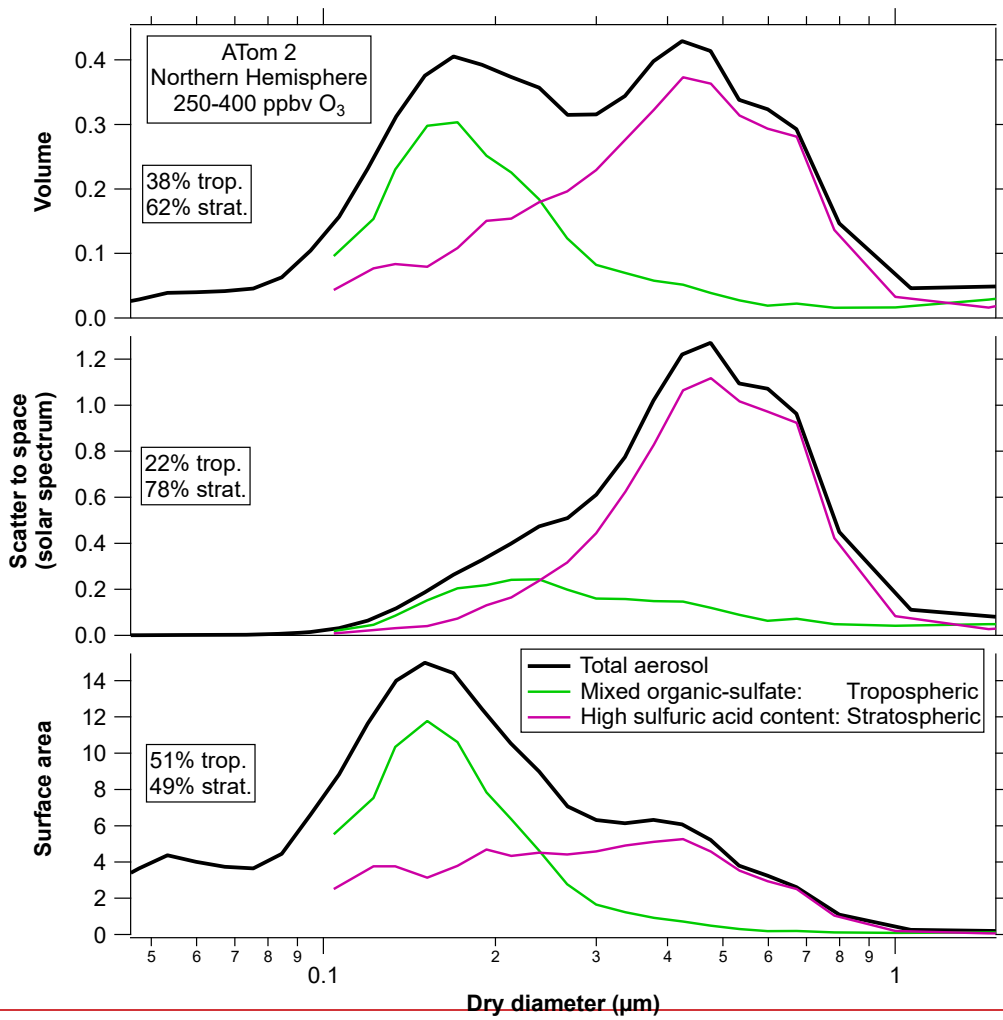
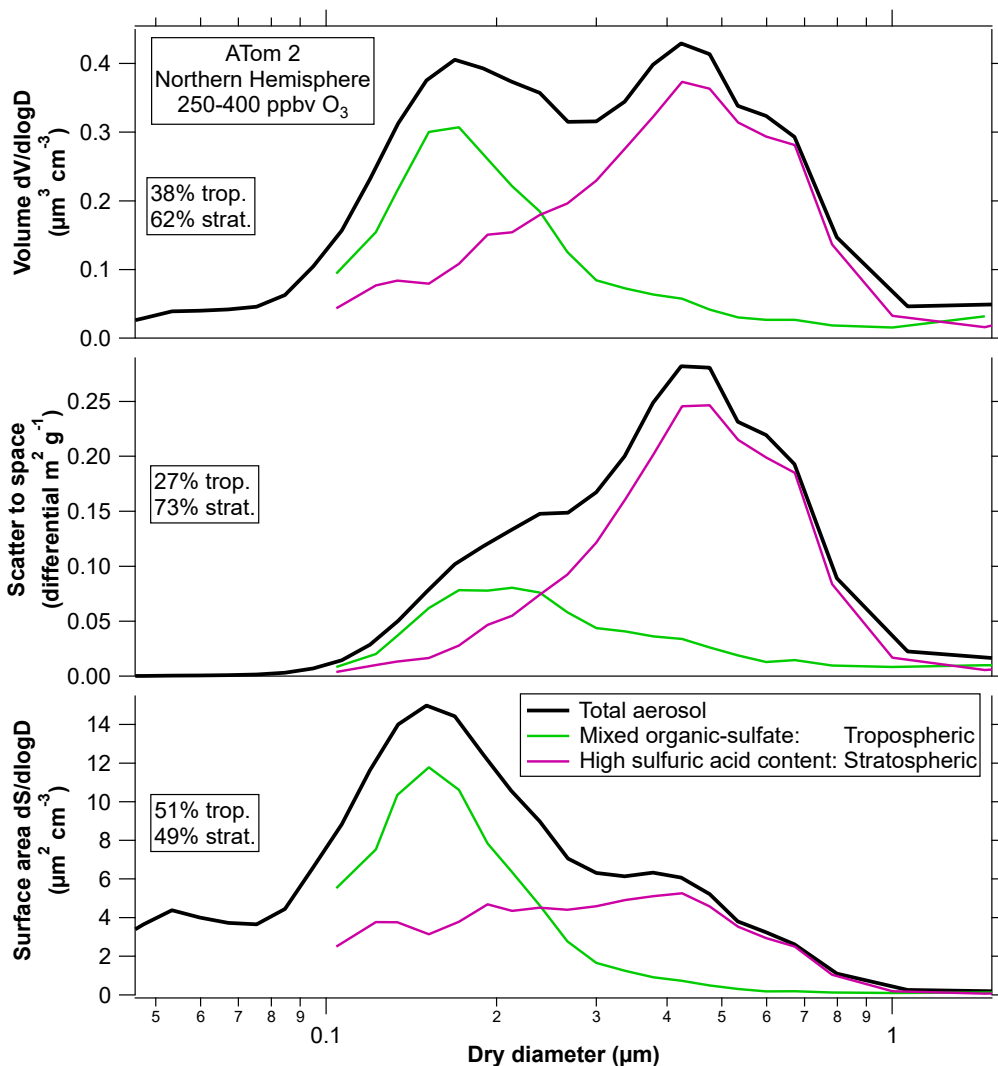


Figure 67. The mass scattering efficiency of particles with a refractive index similar to sulfuric acid. Shown are the total scattered light, the light scattered away from the Earth to space, and their ratio. Larger particles have more forward scattering and therefore a smaller fraction scattered to space. The lowest curve shows the climate impact after an approximate correction for warming due to infrared absorption. Particles smaller than about 0.1  $\mu\text{m}$  or larger than about 4  $\mu\text{m}$  warm the Earth (Lacis et al., 1992). Tropospheric particles containing ammonium sulfate or organic material with higher refractive indices would shift the curves to slightly larger diameters.





**Figure 78.** The composition-resolved size distribution from Figure 1 for the ATom2 Northern Hemisphere weighted by volume, light scattered to outer space, and surface area. The two sulfuric acid categories (with and without meteoric metals) have been combined. Percentages refer to sizes above 0.1  $\mu\text{m}$  only. ATom2 is chosen as an example; Figure S3 S4 shows the percentage contributions of various particle types to these processes for other deployments and for the Southern Hemisphere. Scaling to net thermal infrared heating gives nearly identical relative contributions as volume in the top panel. Scaling to sedimentation rate gives a similar shape to the middle panel.

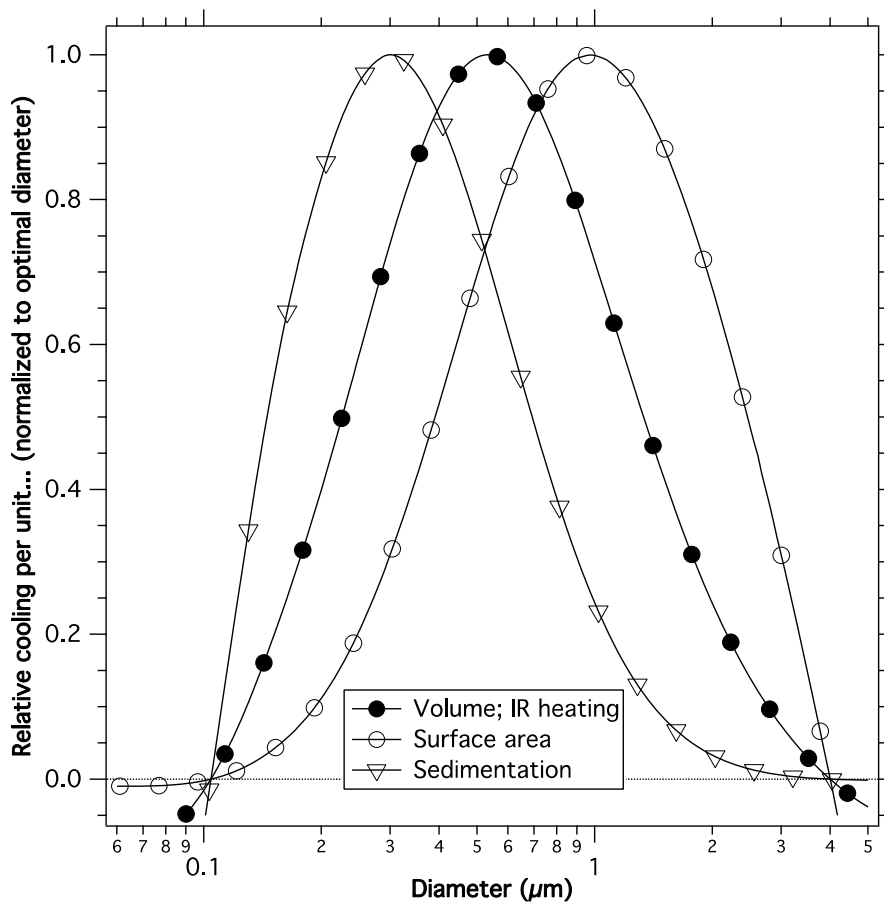


Figure 89. Calculated net cooling per unit volume, surface area, and sedimentation velocity. Calculations of sunlight reflected to outer space are for a single size aerosol averaged over the solar spectrum and zenith angles characteristic of the sunlit Earth. The wavelength averaging eliminates the oscillations from Mie scattering. Particles of about 0.3 μm diameter give the most cooling per unit sedimentation rate and particles of about 0.5 μm diameter give the most cooling per unit mass. Each curve is normalized so the maximum point is unity. For example, 0.2 μm particles have about one tenth the net cooling per unit surface area as 1 μm diameter particles ~~have about 10 times the net cooling per unit surface area than do particles with about 0.2 μm diameter.~~

	Total Volume		Meteoric-sulfuric			Sulfuric			Organic-sulfate			Other
	6 nm-1.7 µm	0.1-1.7 µm	Volume	Fit diam.	Fit σ	Volume	Fit diam.	Fit σ	Volume	Fit diam.	Fit σ	Volume
	µm <sup>3</sup> cm <sup>-3</sup>	µm <sup>3</sup> cm <sup>-3</sup>	µm <sup>3</sup> cm <sup>-3</sup>	µm		µm <sup>3</sup> cm <sup>-3</sup>	µm		µm <sup>3</sup> cm <sup>-3</sup>	µm		µm <sup>3</sup> cm <sup>-3</sup>
ATom1 NH	0.35	0.30	0.046	0.37	1.35	0.103	0.29	1.55	0.150	0.13	1.89	0.004
ATom2 NH	0.35	0.33	0.083	0.46	1.44	0.115	0.35	2.14	0.124	0.17	1.45	0.005
ATom3 NH	0.78	0.76	0.150	0.54	1.51	0.160	0.50	1.74	0.436	<i>0.39</i>	<i>2.89</i>	0.024
Atom4 NH	0.68	0.64	0.066	0.44	1.49	0.175	<i>0.22</i>	<i>2.15</i>	0.292	<i>0.30</i>	<i>2.04</i>	0.107
ATom1 SH	0.49	0.48	0.101	0.39	1.27	0.330	0.40	1.29	0.050	0.27	1.43	0.011
ATom2 SH	0.32	0.31	0.110	0.40	1.39	0.147	0.39	1.64	0.052	0.12	1.80	0.001
ATom3 SH	0.28	0.26	0.056	0.40	1.29	0.088	0.35	1.73	0.120	0.19	1.93	0.003
ATom4 SH	0.19	0.17	0.040	0.36	1.47	0.090	0.16	2.30	0.033	<i>0.20</i>	<i>2.94</i>	0.004
ATom2 NH high	0.40	0.39	0.156	0.42	1.36	0.122	0.40	1.64	0.112	0.18	1.36	0.005
ATom4 NH high	0.37	0.35	0.132	0.41	1.59	0.104	0.31	2.57	0.093	<i>0.38</i>	<i>3.64</i>	0.021

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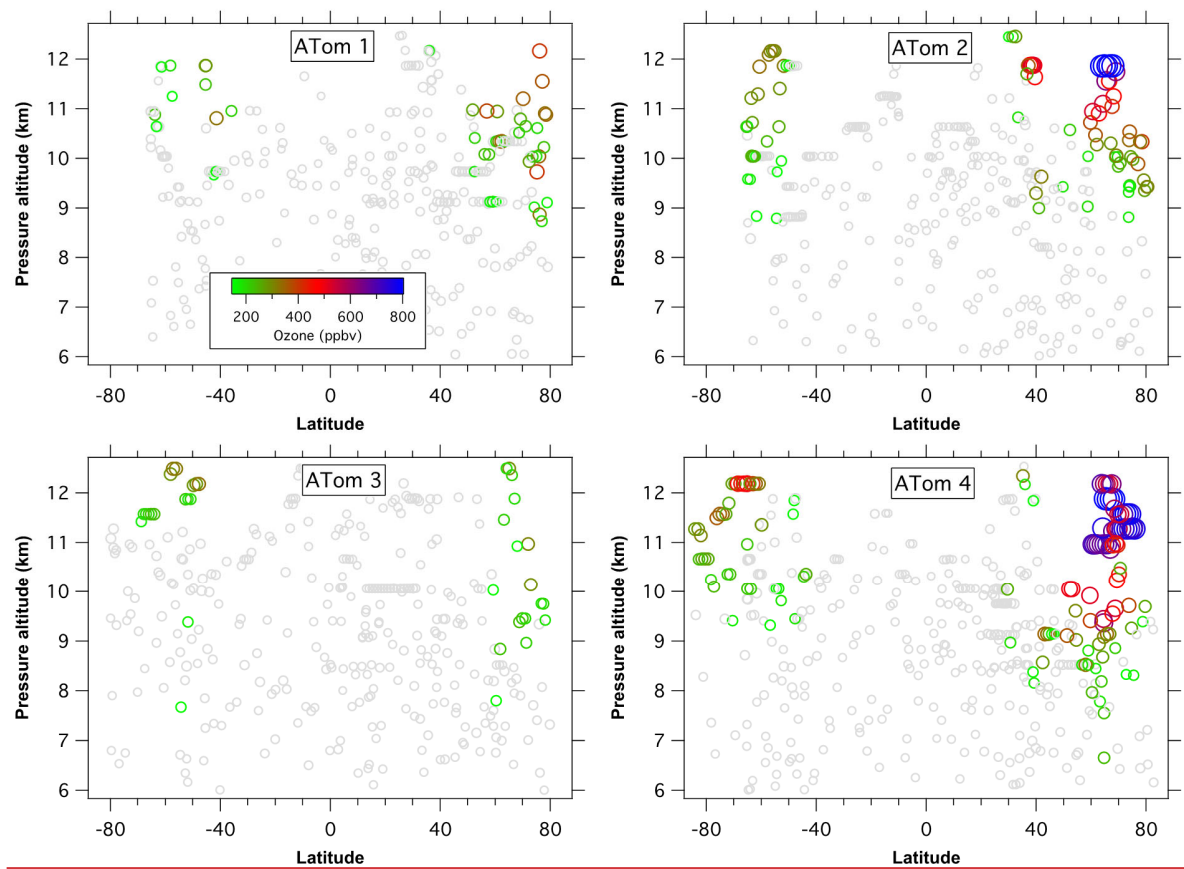
**Table 1. Total volumes and volumes and log-normal fit parameters for individual types of particles. Lognormal fits were performed for particle volume so the fit diameter is for the volume distribution. The volumes and fits are for the size range 0.1 to 1.7 µm diameter except as noted (first column). Data are for 250-400 ppbv ozone except the last two rows are for 500-850 ppbv. Volumes are per standard cm<sup>3</sup>. Fits are to the equation  $V = a \cdot \exp(-(\ln D - \ln D_{fit})^2 / (2\sigma^2))$  where  $D$  is the diameter and  $a$  is a scaling factor. Each fit is for a specific component. Fits were not attempted for “other” particles (mostly dust) because they did not generally show a defined mode. Italics indicate a poor fit to a single lognormal shape.**

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**Supplemental figures for**

Radiative and chemical implications of the size and composition of aerosol particles in the existing or modified global stratosphere

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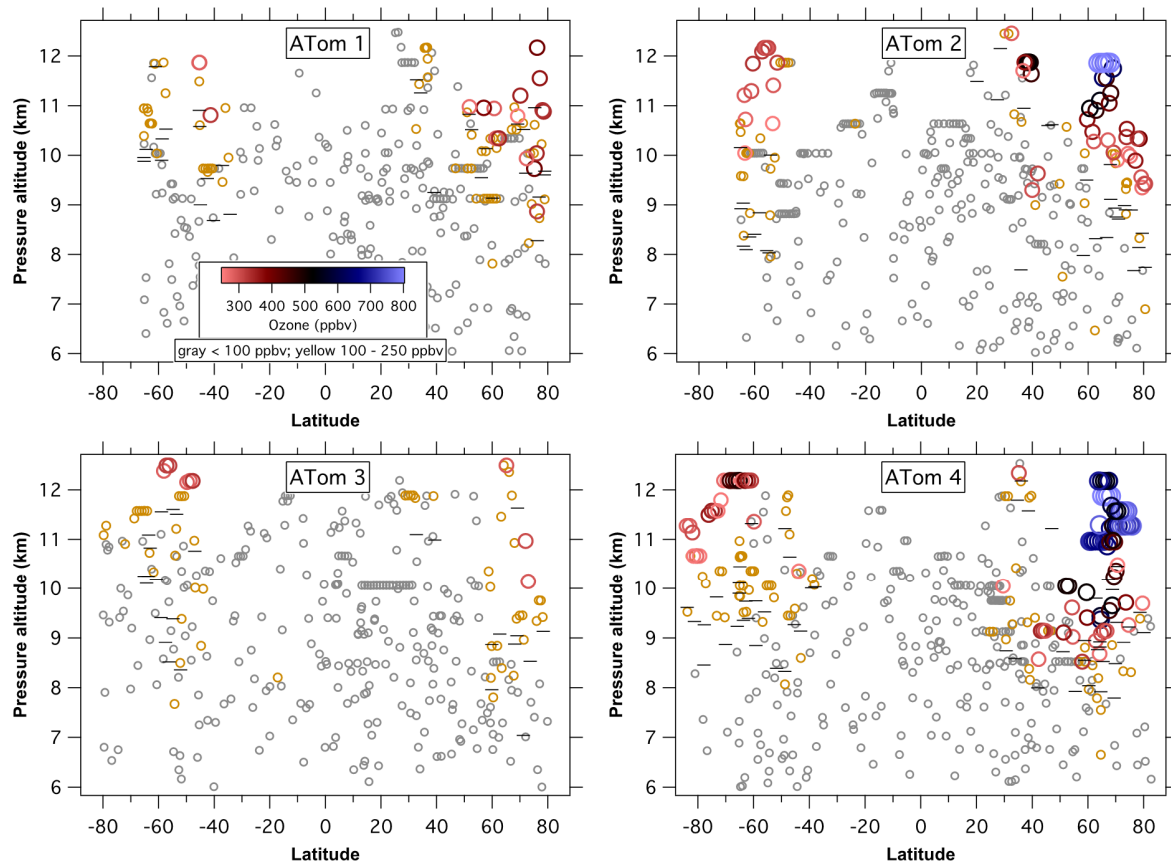
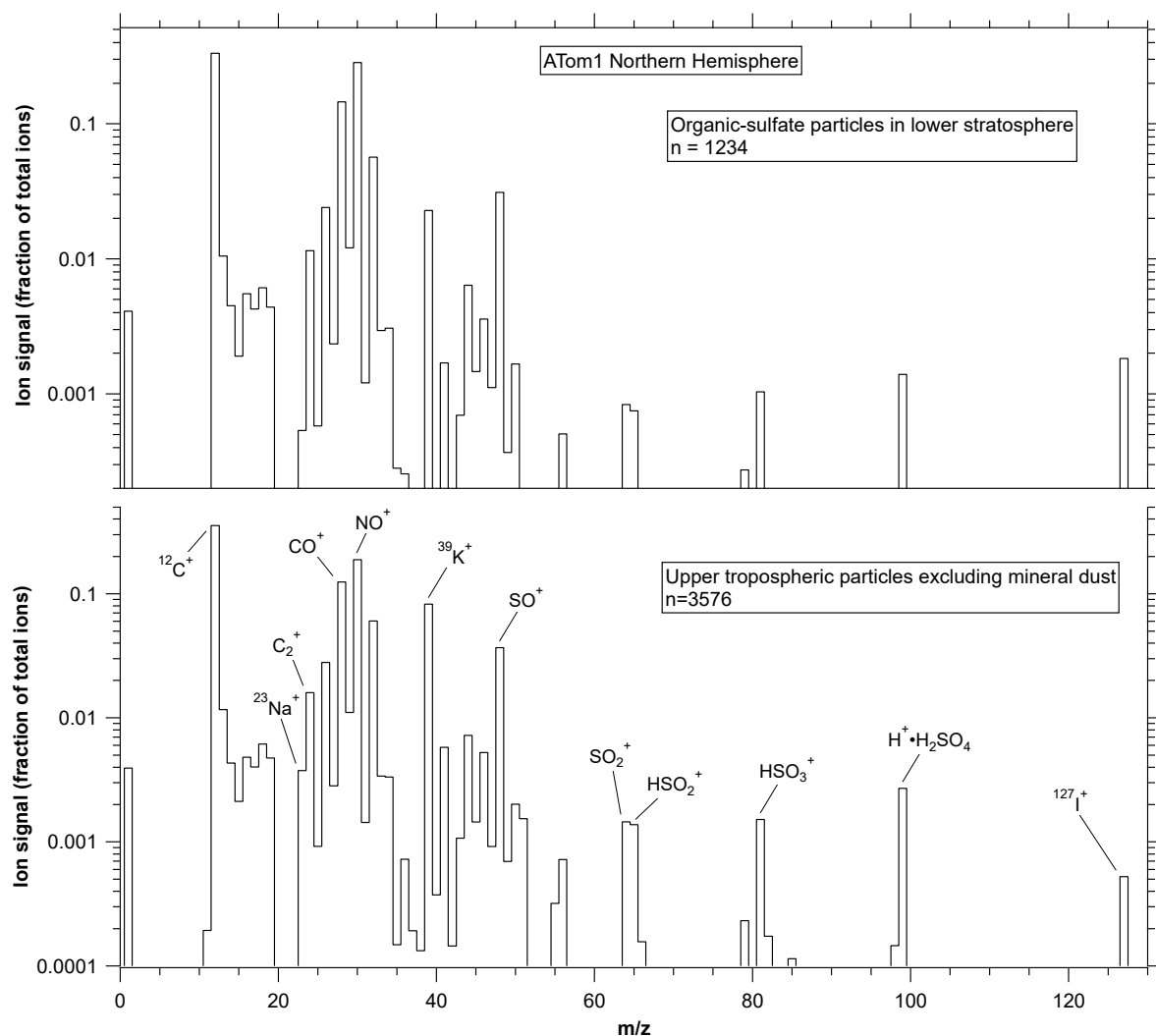


Figure S1. Locations of air with more than 1050 or 250 ppbv ozone sampled during the ATom deployments. Each point represents 500 seconds of cloud-free air when PALMS was sampling. Small horizontal bars indicate locations of the lapse rate thermal tropopause when the DC8 crossed it during an ascent or descent. Other tropopause crossings were in level flight, where the thermal tropopause is hard to diagnose, and are not shown.



**Figure S2. A comparison of the average positive ion mass spectra of organic-sulfate particles smaller than 0.35  $\mu\text{m}$  in the lower stratosphere (ozone > 250 ppbv) with that of particles of similar size in the upper troposphere. Some major peaks with fairly unique identifications are labeled. Other deployments and locations showed similar results. The spectra strongly indicate that the organic-sulfate particles in the lower stratosphere originated in the troposphere.**

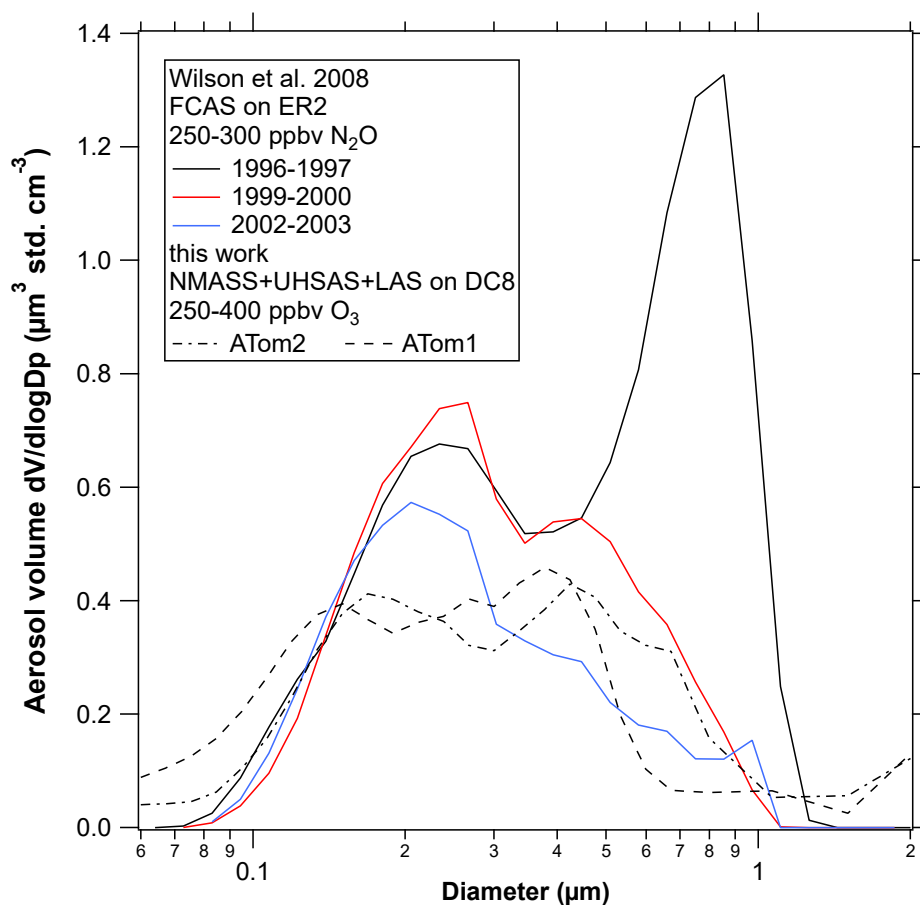
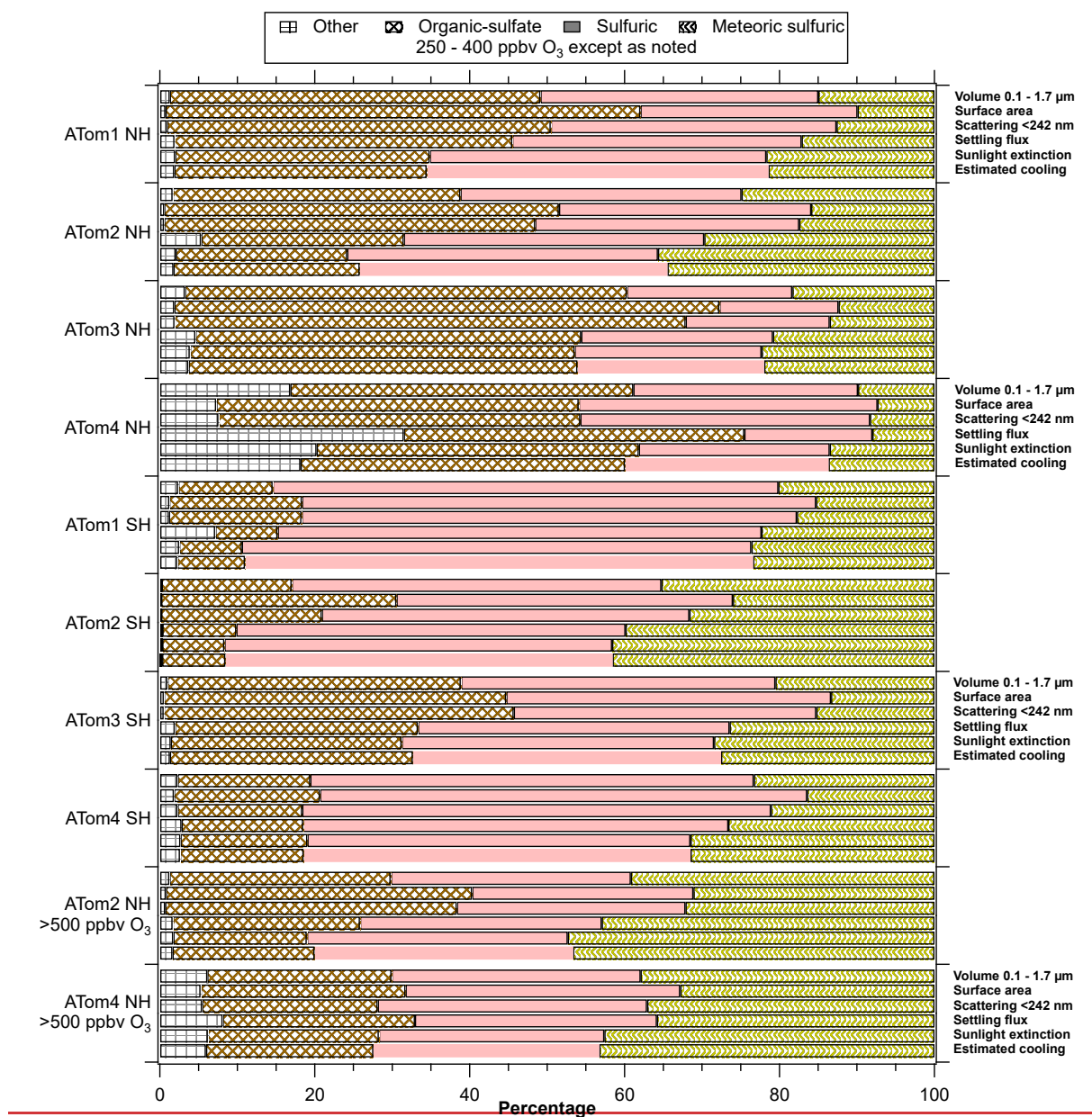
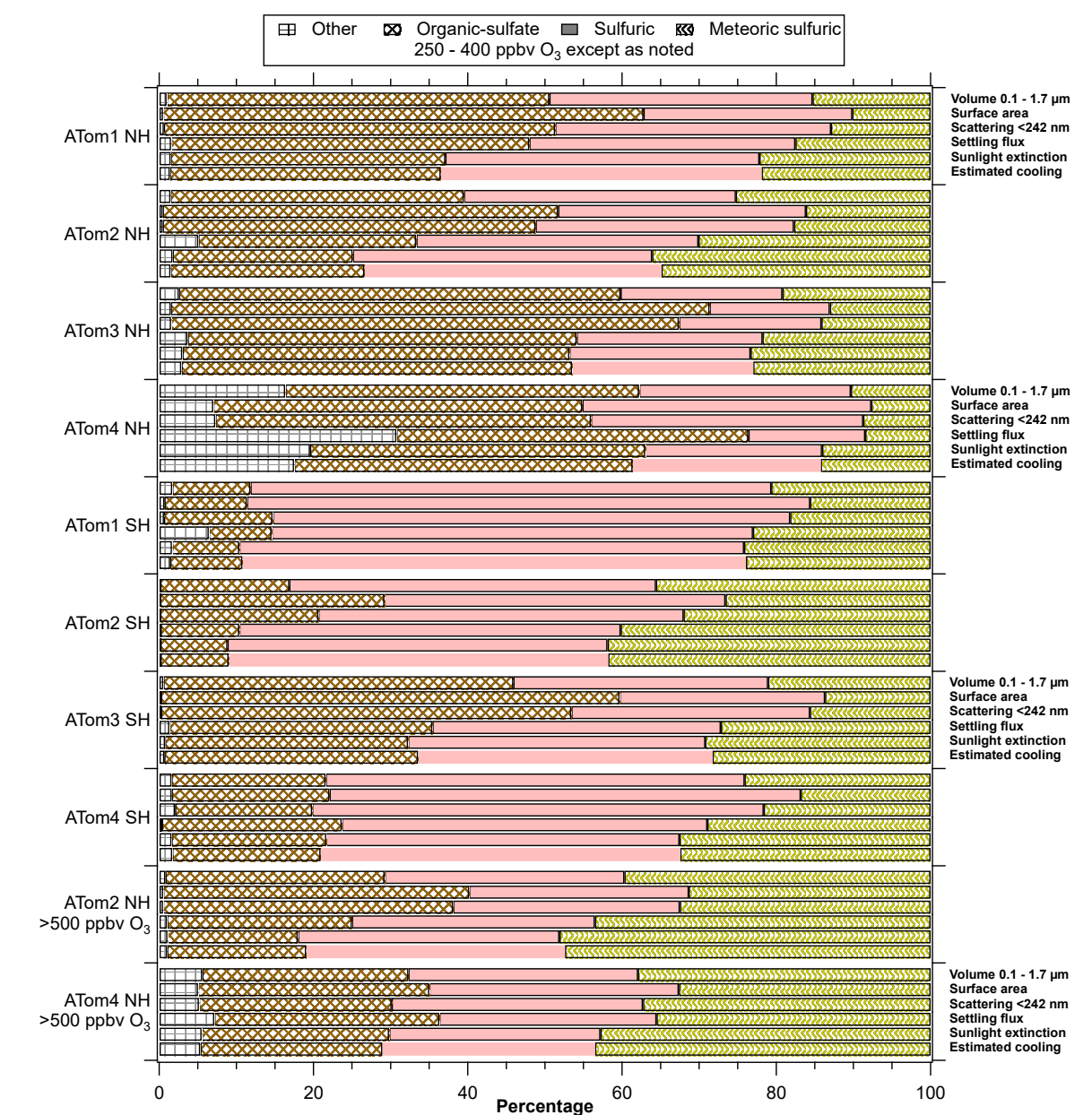
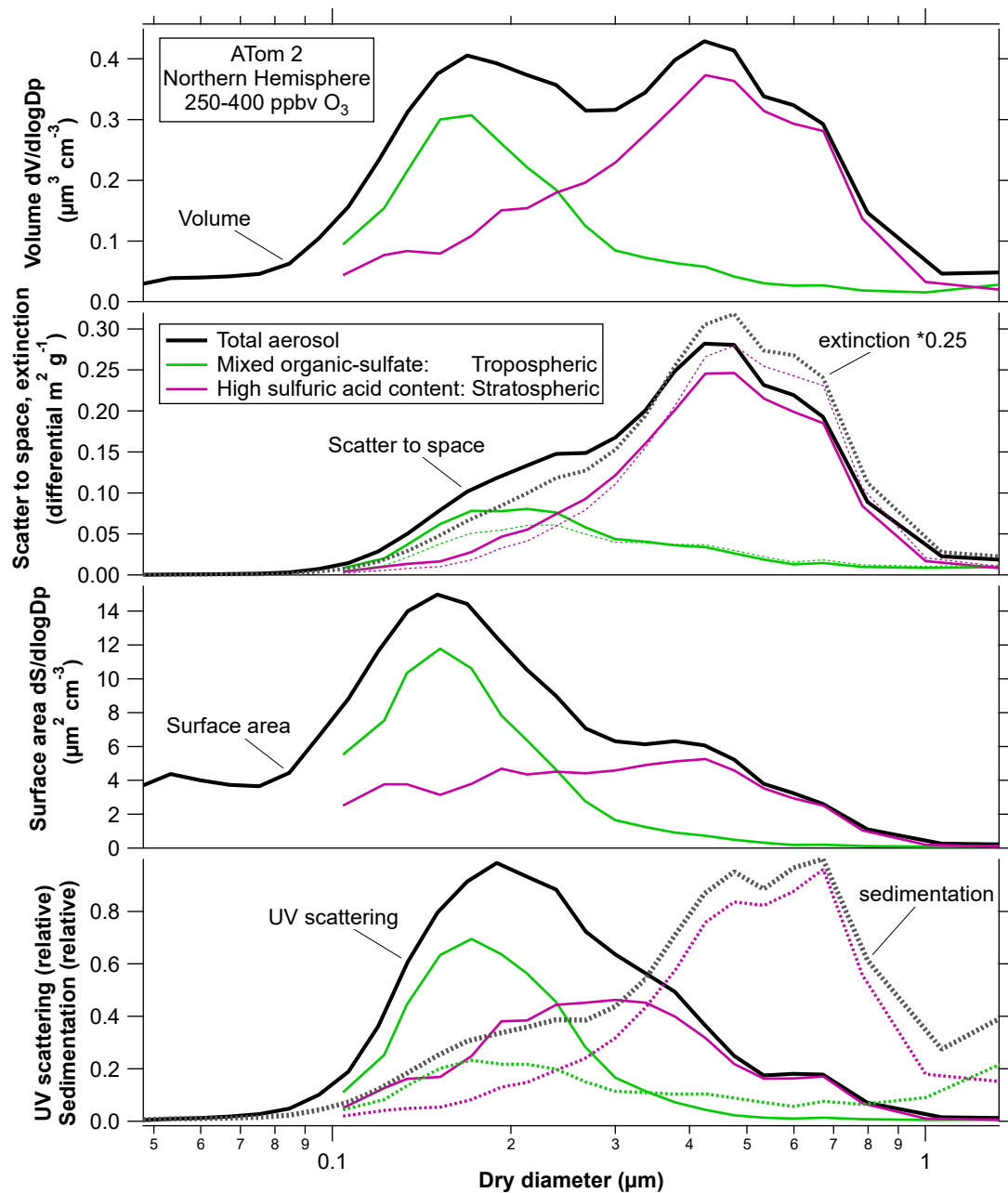


Figure S2S3. A comparison of this work with size distributions in Wilson et al. (2008). These are absolute concentrations instead of the normalized concentrations in the bottom panel of Figure 2 in Wilson et al. (2008). All data are for the Northern Hemisphere. 250 to 400 ppbv of ozone corresponds to roughly 275 to 300 ppbv of N<sub>2</sub>O in the late 1990s. The 1996 Wilson et al. data show the aerosol still influenced by the Mt. Pinatubo eruption after the largest particles had sedimented out.





**Figure S3S4.** Percentages of various effects of dry aerosols in the lowermost stratosphere by type of particle. All calculations are for 0.1 to 1.7 μm diameter. Fractional abundances were extended to 0.1 μm when the data ended slightly above that (see Figure 1). Scattering <242 nm means scattering in the 210-242 nm range important for O<sub>2</sub> photolysis. Settling flux is for 20 km altitude pressure. The settling percentages are not very sensitive to pressure. Estimated cooling is backscatter of sunlight to outer space minus an offset independent of size to estimate the warming due to infrared absorption. Infrared heating for sulfuric acid was used for all particle types because there are insufficient infrared spectral data for the other particle types.



**Figure S5.** Similar to Figure 8 in the paper, data from ATom2 in the Northern Hemisphere. Size distributions weighted by extinction, UV scattering, and sedimentation are shown in addition to the volume, scatter to space, and surface area shown in Figure 8.

250 to 400 ppbv ozone						
Date	Latitude range		Longitude range		# positive	# negative
<a href="#">20160801</a>	<a href="#">52.5</a>	<a href="#">77.1</a>	<a href="#">148.3 W</a>	<a href="#">123.0 W</a>	<a href="#">2416</a>	<a href="#">2531</a>
<a href="#">20160808</a>	<a href="#">-45.5</a>	<a href="#">-36.1</a>	<a href="#">178.3 W</a>	<a href="#">178.2E</a>	<a href="#">1462</a>	<a href="#">1168</a>
<a href="#">20160812</a>	<a href="#">-63.9</a>	<a href="#">-57.5</a>	<a href="#">145.1 W</a>	<a href="#">79.5 W</a>	<a href="#">149</a>	<a href="#">89</a>
<a href="#">20160820</a>	<a href="#">69.2</a>	<a href="#">78.8</a>	<a href="#">77.0 W</a>	<a href="#">49.7 W</a>	<a href="#">2139</a>	<a href="#">1816</a>
<a href="#">20160822</a>	<a href="#">51.7</a>	<a href="#">61.0</a>	<a href="#">85.8 W</a>	<a href="#">70.4 W</a>	<a href="#">957</a>	<a href="#">850</a>
<a href="#">20170126</a>	<a href="#">31.4</a>	<a href="#">33.3</a>	<a href="#">121.0 W</a>	<a href="#">119.7 W</a>	<a href="#">380</a>	<a href="#">240</a>
<a href="#">20170129</a>	<a href="#">49.9</a>	<a href="#">69.7</a>	<a href="#">155.2 W</a>	<a href="#">140.7 W</a>	<a href="#">1607</a>	<a href="#">1140</a>
<a href="#">20170205</a>	<a href="#">-54.0</a>	<a href="#">-51.1</a>	<a href="#">165.8 E</a>	<a href="#">168.9 E</a>	<a href="#">842</a>	<a href="#">613</a>
<a href="#">20170210</a>	<a href="#">-65.3</a>	<a href="#">-54.4</a>	<a href="#">149.8 W</a>	<a href="#">72.1 W</a>	<a href="#">3021</a>	<a href="#">1606</a>
<a href="#">20170218</a>	<a href="#">70.2</a>	<a href="#">75.4</a>	<a href="#">63.7 W</a>	<a href="#">51.6 W</a>	<a href="#">1793</a>	<a href="#">1138</a>
<a href="#">20170219</a>	<a href="#">66.9</a>	<a href="#">80.5</a>	<a href="#">148.1 W</a>	<a href="#">69.3 W</a>	<a href="#">5920</a>	<a href="#">3828</a>
<a href="#">20170221</a>	<a href="#">36.4</a>	<a href="#">52.6</a>	<a href="#">142.0 W</a>	<a href="#">121.8 W</a>	<a href="#">1584</a>	<a href="#">1020</a>
<a href="#">20171008</a>	<a href="#">-52.8</a>	<a href="#">-47.2</a>	<a href="#">166.3 E</a>	<a href="#">168.5 E</a>	<a href="#">1707</a>	<a href="#">1291</a>
<a href="#">20171014</a>	<a href="#">-63.9</a>	<a href="#">-55.2</a>	<a href="#">69.3 W</a>	<a href="#">63.7 W</a>	<a href="#">70</a>	<a href="#">45</a>
<a href="#">20171025</a>	<a href="#">63.0</a>	<a href="#">78.3</a>	<a href="#">151.0 W</a>	<a href="#">77.3 W</a>	<a href="#">2580</a>	<a href="#">1757</a>
<a href="#">20180427</a>	<a href="#">42.1</a>	<a href="#">70.3</a>	<a href="#">154.3 W</a>	<a href="#">132.6 W</a>	<a href="#">3498</a>	<a href="#">2701</a>
<a href="#">20180506</a>	<a href="#">-65.1</a>	<a href="#">-65.0</a>	<a href="#">114.0 W</a>	<a href="#">112.1 W</a>	<a href="#">58</a>	<a href="#">177</a>
<a href="#">20180509</a>	<a href="#">-86.0</a>	<a href="#">-56.4</a>	<a href="#">69.0 W</a>	<a href="#">40.7 W</a>	<a href="#">6336</a>	<a href="#">3843</a>
<a href="#">20180512</a>	<a href="#">-44.2</a>	<a href="#">-43.4</a>	<a href="#">49.8 W</a>	<a href="#">48.5 W</a>	<a href="#">517</a>	<a href="#">275</a>
<a href="#">20180514</a>	<a href="#">29.2</a>	<a href="#">35.7</a>	<a href="#">27.3 W</a>	<a href="#">27.0 W</a>	<a href="#">779</a>	<a href="#">439</a>
<a href="#">20180517</a>	<a href="#">61.6</a>	<a href="#">68.9</a>	<a href="#">57.6 W</a>	<a href="#">28.5 W</a>	<a href="#">1596</a>	<a href="#">856</a>
<a href="#">20180519</a>	<a href="#">57.5</a>	<a href="#">80.0</a>	<a href="#">148.9 W</a>	<a href="#">80.5 W</a>	<a href="#">4705</a>	<a href="#">3544</a>
<a href="#">20180521</a>	<a href="#">38.3</a>	<a href="#">60.2</a>	<a href="#">148.3 W</a>	<a href="#">121.9 W</a>	<a href="#">1142</a>	<a href="#">886</a>
> 400 ppbv ozone						
<a href="#">20160801</a>	<a href="#">61.8</a>	<a href="#">75.9</a>	<a href="#">134.0 W</a>	<a href="#">125.5 W</a>	<a href="#">531</a>	<a href="#">417</a>
<a href="#">20160820</a>	<a href="#">69.9</a>	<a href="#">78.5</a>	<a href="#">75.8 W</a>	<a href="#">54.1 W</a>	<a href="#">341</a>	<a href="#">201</a>
<a href="#">20160822</a>	<a href="#">51.8</a>	<a href="#">56.7</a>	<a href="#">85.7 W</a>	<a href="#">78.8 W</a>	<a href="#">227</a>	<a href="#">99</a>
<a href="#">20170129</a>	<a href="#">59.7</a>	<a href="#">69.5</a>	<a href="#">155.0 W</a>	<a href="#">146.6 W</a>	<a href="#">6152</a>	<a href="#">3619</a>
<a href="#">20170219</a>	<a href="#">67.6</a>	<a href="#">78.9</a>	<a href="#">148.1 W</a>	<a href="#">112.4 W</a>	<a href="#">1320</a>	<a href="#">790</a>
<a href="#">20170221</a>	<a href="#">37.3</a>	<a href="#">39.8</a>	<a href="#">131.7 W</a>	<a href="#">124.0 W</a>	<a href="#">1027</a>	<a href="#">676</a>
<a href="#">20180427</a>	<a href="#">54.5</a>	<a href="#">70.2</a>	<a href="#">154.1 W</a>	<a href="#">142.4 W</a>	<a href="#">7531</a>	<a href="#">3901</a>
<a href="#">20180509</a>	<a href="#">-76.0</a>	<a href="#">-62.4</a>	<a href="#">68.5 W</a>	<a href="#">63.3 W</a>	<a href="#">641</a>	<a href="#">416</a>
<a href="#">20180517</a>	<a href="#">64.4</a>	<a href="#">75.9</a>	<a href="#">66.9 W</a>	<a href="#">42.6 W</a>	<a href="#">4892</a>	<a href="#">3173</a>
<a href="#">20180519</a>	<a href="#">57.3</a>	<a href="#">73.9</a>	<a href="#">148.9 W</a>	<a href="#">80.2 W</a>	<a href="#">7686</a>	<a href="#">4503</a>
<a href="#">20180521</a>	<a href="#">51.0</a>	<a href="#">60.1</a>	<a href="#">148.0 W</a>	<a href="#">141.0 W</a>	<a href="#">2618</a>	<a href="#">1630</a>

Table S1. The number of positive and negative ion mass spectra acquired for two ranges of ozone characteristic of the stratosphere, along with the latitude and longitude of those spectra. Data are above 7 km, out of cloud, and for particles with an aerodynamic diameter acquired so they can be mapped to the size distributions. There are more positive ion spectra because PALMS was programmed to spend more time in positive ion mode. For reference, the total number of positive mass spectra above 7 km (both upper troposphere and stratosphere) was ATom1: 174451; ATom2: 160754; ATom3: 180033; ATom4: 294232.