Significant contrasts in aerosol acidity between China and the Uniteds States

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Abstract. Aerosol acidity governs several key processes in aerosol physics and chemistry, thus affecting aerosol mass and 15 composition, and ultimately the climate and human health. Previous studies have reported the aerosol pH values separately in China and the United States (US), implying a different aerosol acidity between these two countries. However, there is a debate about whether mass concentration or chemical composition plays-is athe more important role in drivinger such aof differences in aerosol acidity. A full picture of the pH difference and the underlying mechanisms responsible for the pH difference it are is hindered not fully understood, limited by by the scarcity of simultaneous measurements of aerosol-particle 20 composition and gaseous species expected and gas species, especially in China. Here we conduct a comprehensive assessment of the aerosol acidity in China and the USUnited States, using extended ground-level measurements and regional chemical transport model simulations. We show that aerosol in China isis significantly less acidic than that in the-USUnited States, with pH values 1–2 units higher. Based on a proposed multivariable Taylor Series method and a series of sensitivity tests, we identify several-major factors leading to the pH difference. Compared to the -USUnited-States, China is associated 25 withhas much higher aerosol mass concentrations aerosol (gas + particle, by a factor of 8.4 on average) and a higher fraction of s in China are generally in more total total ammonia ($TNH_3 = NH_4$ + gas + particle + NH_3) in the aerosol composition rich conditionss where particle phase ammonium (NH_4^+) concentrations are adequate enough to nearly neutralize major acidic inorganic anions such as sulfate, nitrate, and chloride, leading to a higher aerosol pH. Our assessment shows that such differences in mass concentrations and chemical composition play equally important roles in driving the aerosol pH 30 difference between China and the-USUnited States. Therefore, both the facts that China is more polluted than the US and is rich in ammonia together explain the aerosol pH difference. Higher relative availability of the stronger acidic component, sulfate, compared with the weaker acidic component, total nitrate (TNO₃=NO₃⁻+HNO₃), also contributes to the lower aerosol pH in the United States. As a response to higher aerosol pH, the higher nitrate to sulfate molar ratios in China indicates a

nitrate-rich condition, further leading to higher aerosol water uptake which will continually promote nitrate aerosol

35 formation. Considering the historical emissions trends, especially the recent faster decline in SO₂-emissions than NO_{*} emissions in China, the nitrate to sulfate molar ratios in China is expected to further increase, and the difference in aerosol acidity between these two countries is expected to continue the difference in aerosol acidity between these two countries is expected to continue as SO₂ and NO₂ emissions are further controlled. The differences in aerosol acidity highlighted in the present study *imply-implies* potential differences in formation mechanisms, physicochemical properties, 40 and toxicity of aerosol particles between in China and the United States thesese two countries.

1 Introduction

As an intrinsic aerosol property, aerosol acidity (usually measured characterized by aerosol pH) plays an important role in a variety of aerosol physical and chemical processes (Pye et al., 2020). Aerosol acidity can modulate aerosol mass by 45 controlling the gas-particle partitioning of volatile and semi-volatile acids (such as HCl-Cl⁻ and HNO₃-NO₃⁻) (Guo et al., 2016) and <u>can influenceing</u> production rates of secondary aerosol from through heterogeneous pathways (Jang et al., 2002; Surratt et al., 2010; Pathak et al., 2011). HAcidity also affects aerosol optical properties via proton dissociation of organic functional groups (Mo et al., 2017) and the morphology or phase state of organic aerosols (Losey et al., 2016; Losey et al., 2018). Recent evidence links aerosol acidity to aerosol toxicity and health outcomes. For example, Hhighly acidic aerosols, 50 for example, cause greater dissolveution more of metals which can generate reactive oxygen species in vivo (Fang et al., 2017). High aerosol acidity is associated with increased risks of respiratory disease and cancer (Kleinman et al., 1989; Gwvnn et al., 2000: Behera et al., 2015).

Due to the difficulties in directly measurementsing of aerosol pH (Jang et al., 2002; Li and Jang, 2012), thermodynamic models, including ISORROPIA-II (Fountoukis and Nenes, 2007), E-AIM (Clegg et al., 1998), and SCAPE2 (Kim and 55 Seinfeld, 1995), have been widely used to calculate aerosol pH from-based on measured gaseous and aerosol-particle composition and meteorological data such as relative humidity (RH) and temperature. A large number of Manyultiple studies suggest that these models are capable of can reproduce ing the partitioning of semi-volatile species including HNO₃-NO₃⁻ and NH₄⁺-NH₃, which are sensitive to aerosol pH (Guo et al., 2015; Hennigan et al., 2015; Guo et al., 2016).

Field Analyses of field observations in different regions of the United States (US) United States have indicated that the 60 aerosol acidity was is typically high. For example, Weber et al. (2016) (Weber et al., 2016) showed that aerosol pH in the Southeastern United StatesUS was buffered to be nearly constant consistently in the range of -0-2 despite a substantial sulfate reductions over the past 15 years, and this the same trend may be applicable to many other regions. Studies in the <u>Nn</u>ortheastern <u>United StatesUS</u> and California also indicated found -a highly acidic aerosols with mean pH values of 0.8 and 1.9, respectively (Guo et al., 2017a). The aAerosol pH in the Mmidwestern United States US was typically higher than other 65 areas, with anthe average aerosol pH was of 3.8 (Lawal et al., 2018). Studies in China, on the other hand, showed have <u>found</u> generally higher levels of aerosol pH, <u>indicating a lower aerosol acidity</u>. Several studies in the heavily polluted North China Plain (NCP) region, reported average pH <u>values betweenof</u> 3.5–5.2 (Shi et al., 2017; Ding et al., 2019; Shi et al., 2019; Song et al., 2019; Wang et al., 2020a). Xi'an, a city <u>located in Nn</u>orthwest China, had aerosol pH values up to 5 (Wang et al., 2016; Guo et al., 2017b). Some sites in <u>Ss</u>outheast China <u>showed ahad</u> lower aerosol pH, such as the site in Guangzhou with anwhich had an average <u>value</u> of 2.3 (Jia et al., 2020). A comprehensive, nationwide comparison of aerosol pH between China and the <u>US United States can give usprovide</u> a better understanding of the <u>driving</u> factors <u>of driving</u> aerosol pH and its

effect <u>ion</u> aerosol formation mechanisms and properties (Pathak et al., 2009; Guo et al., 2017a; Wang et al., 2020a). However, such comparisons are still scarce (Guo et al., 2017b; Nenes et al., 2020; Zheng et al., 2020), primarily <u>limited</u> <u>bybecause of</u> a lack of extensive simultaneous measurements of aerosol composition and semi-volatile gaseous compounds in China.

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In this study, we compared the aerosol<u>mass concentrations, chemical</u> composition, and acidity between China and the <u>US</u> <u>United States</u> based on one-year <u>large scale</u>-measurements from 34 ground monitoring sites in the <u>United StatesUS</u> and 16 sites in China<u>(mostly clustered in the NCP)</u>. In order to extend the spatial coverage to nationwide scales, we employed the Community Multiscale Air Quality (CMAQ) model to simulate the concentrations of gaseous and <u>aerosol-particle</u> species which were used to calculate aerosol pH <u>in-across</u> both countries. We proposed a new method to identify the factors driving <u>factors leading to-the pH difference between these two countries then carried out a series of sensitivity tests to identify and discussed the causes and implications<u>effects associated withof</u> of the pH difference.</u>

2 Data collection and method

2.1 Observational data

- Measurements of Ggaseous species (including HNO₃, NH₃ and HCl) and, aerosol-particle_components (including PM_{2,5} components of SO₄²⁻, NO₃⁻, NH₄⁺, and Cl⁻), and nonvolatile cations (NVCs)) (including Na⁺, Mg²⁺, K⁺, and Ca²⁺) collected from monitoring networks in China and the United StatesUS and are used for analysis and comparison in this study. NVCs includeconsidered are Na⁺, Mg²⁺, K⁺, and Ca²⁺. The names and locations of the monitoring sites can be found in Tables S1 and S2. The sum of total observed aerosol ionic compounds is defined as water soluble ions (WSI), though it is recognized that not all of the ions are routinely measured, including such as (e.g., trace species and organic ions). We also study the partitioning of semi-volatile species including NH₃--NH₄⁺ and HNO₃-NO₃⁻ because they are sensitive to pH, especially when the partitioning ratios, ε(NH₄⁺) and ε(NO₃⁻), defined as the molar ratio of NH₄⁺ to total ammonia (TNH₃=NH₃+ NH₄⁺) and the molar ratio of NO₃⁻ to total nitrate (TNO₃=HNO₃+NO₃⁻), are around 50% (Guo et al., 2017a; Chen et al., 2019).
- In the <u>United StatesUS</u>, observational data are from <u>the</u>_co-located<u>sites of <u>the</u>_</u>Clean Air Status and Trends Network 95 (CASTNET) (https://www.epa.gov/castnet) and Ammonia Monitoring Network (AMoN) (<u>http://nadp.slh.wisc.edu/amon/</u>) <u>sites</u>. <u>Two sites in</u>_CASTNET and AMoN<u>sites</u> are assumed to be co-located if <u>they are</u> within 1km<u>of one another</u>.
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Observations from co-located sites and are then combined for pH calculation. Weekly ambient concentrations of gases and particles particulate species, including HNO₃, SO₄²⁻, NO₃⁻, NH₄⁺, Cl⁻ and NVCs, are available from CASTNET sites, while biweekly concentrations of NH₃ are available from AMoN sites. In order to To match biweekly data of NH₃ from AMoN to 100 weekly data of other species from CASTNET, we assume concentrations of The same NH₃ in two adjacent consecutive weeks are used for both weeks of the CASTNET samples to be the same. This assumption is expected to have a minor effect on pH prediction study found that a 10 times increase in NH_3 is required to increase pH by one unit (Guo et al., 2017b), and This assumption is also confirmed in the supported in later discussion (Text- \$15ect.3.2.3). HCl data is not available, so we only use particle phase Cl⁻ as total Cl in-for pH calculations. Sensitivity tests either assuming HCl 105 concentrations of a-four times of HCl vs. the Cl⁻ concentrations or useing derived-HCl concentrations derived from CMAQ modeled HCl/Cl⁻ ratios showed little difference in aerosol pH₇ compared to that the pH estimated by using particle phase Cl⁻ as total Cl (Fig-.ure S1). Considering the small reported little change in aerosol pH in the United StatesUS over a long-term period (Weber et al., 2016; Lawal et al., 2018) and the configuration of the chemical transport model which is set up for the year 2011 (see the following section), we use observational data in 2011 to investigate the aerosol pH in the United States US. 110 Only the sites with measurements available for all species were selected infor this study. As a result, tThere are 34 co-located CASTNET and AMoN sites, which are evenly distributed scattered across the contiguous United States US (Fig. S2a). The accuracy for of CASTNET measurements has been assessed through the analysis of reference and continuing calibration verification samples with a criterion of were all in 95%-105% (except for NH4⁺, which is inhas an accuracy of whose accuracy criterion is 90%-110%), while the accuracy of NH₃ data derived from AMoN was in 97% 103%, suggesting a good he high 115 quality of this datasets the data. Detailed information about data quality could be found is available at in the CASTNET

- Quality Assurance Report-Annual 2011 (United States Environmental Protection Agency, 2012b) and Quality Assurance Support for the NADP (National Atmospheric Deposition Program). A previous study demonstrated that the NH₃ concentrations measured by the passive AMoN samplers operated by AMoN are comparable to annular denuder systems (as a reference system) with thea mean relative percent difference of -9% (Puchalski et al., 2015).
- 120 In China, hourly observational data are extracted from the dData-sharing platform operated by the Comprehensive Observation Network for Air Pollution in Beijing-Tianjin-Hebei and ilts **s**Surrounding Areas (http://123.127.175.60:8765/siteui/index). This collaborative observation network gathers information from is supported by multi source multiple institutions and and provides simultaneous observations of gaseous and aerosol-particle species at individual monitoring sites (Wang et al., 2019). In this study, wWe derive daily average concentrations of gaseous species 125 including NH₃, HNO₃ and HCl and of particle species including NH₄⁺, NO₃⁻, Cl⁻, and NVCs for pH calculation from hourly
- observational data at 16 monitoring sites for use in pH calculation. These monitoring sites are clustered in NCP in-the eastern <u>China (Fig. S2c)</u>. Due to the lack of data quality information, we first processed the data before using them by removing unreasonable data points for quality control by removing unreasonable data. We define a set of valid data of containing all the the-measured components in one day as one case. We first remove the cases with one or more any missing component(s). In
- 130 this step, 2704 out of 5840 cases are removed in this step. We then identify the data points beyond the scope of that is are

more than three median absolute deviations from the median as outliers and remove the cases with any component identified as an outlier. Eventually, 1766 cases remain for subsequent analyses. Although we remove a large number of many cases in this process, the remaining cases cover most of the days in a year and distribute evenly distributed by months (Table S3). For example, data points showing extremely large values for certain species (e.g. [Ca²⁺]>1000µg·m³) are removed. 16 135 monitoring sites with measurements available for all species are selected in this study. These sites cluster in NCP in eastern China (Fig. S2c). It should be noted that the weekly (or longer) duration of the CASTNET samples in the US may lead to biases in the measured concentrations especially for those volatile species such as ammonium nitrate. Sickles et al. (1999) conducted a comprehensive comparison of the measurements using the CASTNET weekly-duration sampling approach with those using a 24-h-duration sampling approach. Both approaches used filter packs. They found that compared to 24-h durations ampling, weekly durations ampling led to low biases of -5%, -5%, and -0.7%, on average, in measured HNO₃, NO₃⁻, and NH_4^+ , respectively, and high biases of 4% and 16%, on average, in SO_4^{2-} and SO_2 , respectively. To evaluate the potential biases pertaining toin the calculated aerosol pH due to the long weekly--duration sampling, we conduct a sensitivity test to adjust the CASTNET-measured concentrations based on these reported average differences between weekly-duration and 24h-duration samples (Sickles et al., 1999) (Results and Discussion).

145 **2.2 Model configuration**

We use the CMAQ version 5.0.2 (United States Environmental Protection Agency, 2014) to simulate gaseous and aerosol particle species concentrations and aerosol pH in China and the United States US. The model domains of the two simulations cover the mainland China and the contiguous United States US with 124×184 and 112×148 horizontal grid cells, respectively. and are bBoth are resolved at the 36-km horizontal resolution and with 13 vertical layers extending to ~16 km above the ground. In both simulations, gas-phase chemistry is modeled with the CB05 chemical mechanism (Yarwood et al., 2005), and the aerosol thermodynamic equilibrium is modeled with ISORROPIA II (Fountoukis and Nenes, 2007).

The meteorological and emission inputs used to drive the China's simulation are adopted from "AiMa", an online operational air quality forecasting system (Lyu et al., 2017; AiMa Forecast, 2017). In the AiMa modeling system, the meteorological data were are generated with the Weather Research and Forecasting (WRF) model (William C. Skamarock 155 2008) driven by the 0.5-degree global weather forecast products produced by the National Centers for Environmental Prediction (NCEP) Global Forecast System (GFS) (Global Forecast System (GFS) Model). The AiMa emission inventory was compiled and derived by integrating a variety of inventories and utilizing various activity data and has been continuously updated since its establishedment (Lyu et al., 2017). The base year of the current AiMa emission inventory is 2017. For the United StatesUS simulation in the United States, we use WRF-modeled meteorological fields downscaled from the North American Regional Reanalysis (NARR) data (Mesinger et al., 2006) as the meteorological input and the 2011 National Emissions Inventory provided by the United StatesUS Environmental Protection Agency as the emission input (United States Environmental Protection Agency, 2012a). The base year of the meteorology and emissions is consistent with the year of the measurements in each country (i.e., 2017 for China and 2011 for the United States US).

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In order to evaluate the model performance against observations, we calculate normalized mean bias (NMB) and normalized root-mean-square errors (NRMSE) to evaluate the spatial variation of pH, species concentrations and partitioning ratios with following equations:-

$$NMB = \frac{\sum_{1}^{N} (c_m - c_o)}{\sum_{1}^{N} c_o}$$
(1)

$$NRMSE = \frac{\sqrt{\frac{\sum_{1}^{N} (c_m - c_0)^2}{N}}}{\frac{N}{C_0}}$$
(2)

where C_m is the <u>CMAQ</u>-modeled value, C_o is the <u>observational-observed</u> value, N is the number of simulation-observation pairs-used in NMB and NRMSE calculations.

2.3 Aerosol pH calculation

In this study, we use the ISORROPIA-II thermodynamic model (Fountoukis and Nenes, 2007) to determine the composition in a K⁺-Ca²⁺-Mg²⁺-NH₄⁺-Na⁺-SO₄²⁻-NO₃⁻-Cl⁻-H₂O aerosol system under equilibrium conditions with gas phase precursors. Aerosol pH is calculated based on H⁺_{air} and liquid water content uptake by inorganic species (LWC_i) from ISORROPIA-II output:

$$pH = -log_{10}(\gamma_{H^+} \cdot H^+_{aq}) = -log_{10}\left(\frac{1000\gamma_{H^+} \cdot H^+_{air}}{LWC}\right)$$
(3)

where γ_{H+} is the activity coefficient of the hydronium ion and which is assumed to be 1 in this study (note that, the binary activity coefficients of ionic pairs, including H⁺, are calculated in ISORROPIA-II), H⁺_{aq} (mol·L⁻¹) is the hydronium ion concentration in the aerosol liquid water, H^+_{air} (µg·m⁻³) is the equilibrium particle hydronium ion concentration per volume air. LWC (µg·m⁻³) in this study only considers the water uptake by inorganic species. includes LWC, and LWC, which means the water uptake by inorganic species and organic species ($\mu g \cdot m^{-3}$) are modeled separately because both organic and

inorganic species are hygroscopic. In this study, we only consider the effect of LWC, since Note that tThe effect of water

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There are two modes in ISORROPIA-II's calculation, i.e. the forward mode and the reverse mode. In the forward mode, the 185 inputs include total concentrations (i.e. gas+aerosolparticle) of TNH₃, TNO₃, TCl (HCl+Cl⁻), SO₄ and NVCs and meteorological parameters (temperature and, relative humidityRH); in the reverse mode, only the aerosol-particle phase of compounds and meteorological parameters are needed (Fountoukis and Nenes, 2007). In this study, the ISORROPIA-II model is run in the forward mode for aerosol in metastable state because the concentrations of both gas and particle species are available and also because the reverse mode has been reported to be more sensitive to measurement errors (Hennigan et 2015; Song et al., 2018).

uptake by organic species LWC_e on aerosol pH has been found to be minor (Guo et al., 2015).

We also-find that there are measurements with unrealistically high Ca^{2+} concentrations (such that Ca^{2+} is more than LWC×0.002, i.e., the solubility of Ca^{2+} in aerosol liquid water). This may be due to the measurement method of Ca^{2+} which needs to uses large amounts of water to dissolve filter-based collected particles. This process will likely dissolve the water-

- insoluble part of Ca^{2+} in aerosols which may cause higher bias of <u>in</u> aerosol Ca^{2+} concentrations. In the existence of aerosol SO₄²⁻, Ca^{2+} precipitates along with SO₄²⁻ as CaSO₄ because of the low solubility (Seinfeld and Pandis, 2006). Including the high Ca^{2+} concentration <u>or not causes aleads to</u> large differences in <u>predictedestimated</u> pH because of <u>the</u> high acidity of SO₄²⁻ (<u>Text S2Sect. 3.2.4</u>). In order to avoid th<u>ise potential</u> bias, we use modified Ca^{2+} concentration for pH calculations while keeping SO₄²⁻ unchanged₇₋ T that is, we use <u>the</u> original Ca^{2+} concentration to calculate aerosol LWC and <u>then</u> use the concentration of Ca^{2+} that can dissolve in the LWC as the modified Ca^{2+} concentration <u>if in cases where the</u> original Ca^{2+} <u>is in</u> excess of exceeds its solubility in the calculated LWC.
- We evaluate the model performance by comparing the gas-particle partitioning of semi-volatile compounds between measured and simulated values such as c(NO₃⁻) and c(NH₄⁺). We compare the directly measured gas-particle partitioning ratios of semi-volatile compounds with the ratios re-partitioned by ISORROPIA-II using measured total (gas+particle) concentrations as inputs. The purpose of this comparison, as conducted in previous studies (Guo et al., 2016; Guo et al., 2017a), is to examine the measurement data quality. This method is effective when the species have substantial fractions in both gas and particle phases (Guo et al., 2017a). The comparison results of ε(NH₄⁺) and ε(NO₃⁻) are shown in Fig. S3. The correlation coefficients and the slopes of linear regression are all close to 1, suggesting good agreement between the <u>-the measured and ISORROPIA-II</u> performs better in the <u>United StatesUS</u> than in China, which may <u>be</u> attribut<u>able</u>, e in partly, to the more evenmore balanced partitioning of the species between gas and particle phase in the <u>United StatesUS</u>.

2.4 Multivariable Taylor Series Method (MTSM)

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In order <u>T</u>to separate the contribution of <u>each-individual</u> components (<u>8-eight</u> species in total, includinge Na⁺, SO₄, TNO₃, TNH₃, TCl, Ca²⁺, K⁺, and Mg²⁺) and meteorological variables (<u>relative humidityRH</u> and temperature) to the pH difference between China and the <u>United StatesUS</u>, we <u>use-propose</u> a multivariable Taylor Series <u>Method-method</u> (MTSM). First, we derive the average conditions (i.e., species concentrations and meteorological conditions) across all the sites in the <u>United StatesUS</u> as the starting point and China as the end point and decompose the contributions of individual compounds to the pH difference based on the following equations:

$$\Delta c_i = c_{i,China} - c_{i,US} \tag{4}$$

$$c_{i,\lambda} \cong c_{i,US} + \Delta c_i \cdot \lambda \tag{5}$$

$$220 \qquad \Delta pH = pH_{US} - pH_{C/ina} = \int_0^1 \left(\sum_{i=1}^8 \frac{\partial pH}{\partial c_{i,\lambda}} \cdot \Delta c_i \right) \cdot d\lambda \cong \sum_{s=1}^{100} \left(\sum_{i=1}^8 \frac{\partial pH}{\partial c_{i,\frac{s}{100}}} \cdot \Delta c_i \right) \cdot 0.01 \tag{6}$$

$$\Delta p H_i \cong \sum_{s=1}^{100} \frac{\partial p H}{\partial c_{i,\frac{s}{100}}} \cdot \Delta c_i \cdot 0.01 \tag{7}$$

where subscript *i* denotes a specific species or meteorological variable; $c_{i,China}$ and $c_{i,US}$ represent the values of concentration species *ithe concentration of compound i* in China and the United StatesUS, respectively; Δc_i is the difference in c_i between

- China and the United StatesUS; $c_{i,\lambda}$ is an intervening c_i between $c_{i,China}$ and $c_{i,US}$ defined by λ , $-\lambda \in [0, 1]$; when λ is 0, $c_{i,\lambda}$ is $c_{i,US}$ when λ is 0; when λ is 1, $c_{i,\lambda}$ is $c_{i,China}$ when λ is 1. In this study we assume negligible interaction between species, therefore the increasing concentration of species *i* will not have the effect of changing the concentration of other species. The pH difference between China and the United StatesUS (i.e., ΔpH) can be expressed as the sum of the partial derivatives of pH with respect to $c_{i,\lambda}$ which is then integrated from $c_{i,US}$ to $c_{i,China}$, as described by Eq. (6). In this study, we take 100 steps with equal intervals to gradually change λ from 0 to 1 (Eq. (6)) and record the partial derivatives of pH with respect to individual $c_{i,\lambda}$, and derive the contributions of all the species and meteorological variables to the pH change at every step. By summing up the contributions of individual variables at all the steps, we characterize the contributions of individual factors
 - components to the overall pH difference (Eq. (7)). Based on the same method, we further quantify the contributions of individual factors to the differences in LWC and H_{air}^+ , respectively, the two variables directly used to calculate aerosol pH (Results and Discussion).

235 3 Results and discussion

3.1 The pH difference between China and the United States US

3.1.1 The pH difference based on observations

- The sensitivity test to adjust the CASTNET-measured concentrations based on the reported average differences between weekly-duration and 24-h-duration samples shows little difference between the unadjusted and adjusted pH values in the US (2.69±0.85 and 2.74±0.83 on average for the unadjusted and adjusted pH, respectively), suggesting that the weekly duration of the CASTNET sampling has little impact on the calculated aerosol pH. Therefore, we proceed with our subsequent analyses using the unadjusted pH. The aerosol pH values calculated based on observational data show a significant difference between China (most observation sites are in NCP) and the United StatesUS. In China (mainly the NCP), the 2017 annual average pH level-at monitoring sites is 4.3-and, rangesing from 3.3 to 5.4 by monitoring sites-with an interquartile
- range of 3.9–4.6. In the <u>contiguous United StatesUS</u>, the 2011 annual average pH level-is 2.6, ranging from 1.9 to 3.9 with an inter-quartile range of 2.2–3.0 (Fig. 1). The t-test shows a <u>statistically</u> significant difference between the two groups (p<0.0001), suggesting that the aerosols are on average more acidic at the monitoring sites in the <u>contiguous United</u> <u>StatesUS</u> than in China (NCP).

The pH difference <u>can-is</u> also <u>be</u>-illustrated by the cumulative distribution function (CDF) curves (Fig. 2, solid lines). The shapes of the CDF curves are similar in these two countries with a slightly steeper slope in the <u>contiguous</u> <u>United StatesUS</u> (Fig_-2a). The pH values, however, are 1–2 units higher in China (<u>NCP</u>) than in the <u>contiguous</u> <u>United StatesUS</u> <u>at</u> across <u>varying</u> levels of cumulative frequencies in the <u>CDF curves</u>. In some cases, <u>the</u> aerosols<u>-acidities</u>_could be completely neutral in China (<u>NCP</u>) (the frequency is 2% for pH \geq 7), while in the <u>contiguous</u> <u>United StatesUS</u>, the pH values in all <u>the</u> cases are below 6.

- Spatially, 14 out of the 16 sampling sites in China are in the NCP (Fig. Figure-Fig. S2c) which is one of the most populous 255 and polluted regions in China (Hu et al., 2014; Cui et al., 2020). Our pH results in this region are consistent with other studies (ranging from 3.5 to 4.6) (Liu et al., 2017; Ding et al., 2019; Ge et al., 2019). The distribution of sampling sites in the United States US, on the other hand, is more evenly distributed spatially. The pH values in the Midwest and California are higher than in other regions like the Southeast, in line with previous studies (Lawal et al., 2018; Chen et al., 2019). Overall, 260 the pH level in the United States US is 1.7 units lower than in-over the NCP of China.

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3.1.2 The pH difference based on model simulations

To solve the uneven spatial address the issue of insufficient spatial coverage of the observational data in China, we conduct simulations using CMAQ, in company-conjunction with the observational data, to further address study the pH difference on a nation-wide scale. We evaluate the model performance by comparing the modeled and observed aerosol pH values (Fig. 3), major aerosol-particle and gaseous species including SO_4^{2-} , NO_3^{-} , NH_4^+ and HNO_3 , NH_3 , and the partitioning ratios including $\varepsilon(NH_4^+)$ and $\varepsilon(NO_3^-)$, at monitoring sites (Figs. S4–S6).

Spatially, the model simulations generally capture the observed variations in pH, species concentrations, and partitioning ratios, although there are some notable biases-occurred (Figs. S4 and S5). In both China (NCP) and the contiguous US, the modeled NH_4^+ , NO_3^- , and NH_3 are biased low, compared to observations, while modeled HNO_3 are biased high, 270 which resultsing in low biases in the predicted $\varepsilon(NO_3^-)$ and $\varepsilon(NH_4^+)$. The modeled SO₄²⁻ in the twoboth countries are all is biased low. For $SO_4^{2^*}$, the model captures the high concentration in the NCP and the eastern US, but it shows low biases in some sites in the southern NCP. This leads to a more negative NMB of the modelled SO4²⁻ in China than in the United States, which can also be seen from Fig S5a& Fig S5i. Low biases are also found for other aerosol components, including NH4⁺, NH3- and NO3⁻, in both countries. Such low biases have been seen in previous studies (Fountoukis et al., 2013; 275 Theobald et al., $2016)_{3}$ which can be and have been attributed to the spatial mismatch between the observations and simulations due to the coarse resolutions of the model grid cells (usually in the range of 20–50 km resolutions) (Shen et al., 2014; Wang et al., 2014a). Smaller NMBs in the US indicated a better performance, compared to China (NCP). Larger differences between observations and simulations in China (NCP) could also be caused by larger measurement uncertainties as the data in China are collected from different monitoring stations operated by individual research institutions (Wang et al., 2802019) and thus lack a unified quality control, compared with data in the US, which comes from national monitoring networks (United States Environmental Protection Agency; National Atmospheric Deposition Program). Due to tThe cooccurrence of the low biases in $\epsilon(NO_3^-)$, which causes lower bias in aerosol pH, and low biases in $\epsilon(NH_4^+)$ and SO_4^{2-} , which cause higher bias in aerosol pH, likely offset each other, there sulting in small biases pertaining to in aerosol pH-likely offset each other, Indeed, and the simulated average pH values in the observation sites $(3.8\pm0.2 \text{ in NCP}, \text{China and } 1.8\pm0.5 \text{ in the})$ 285 contiguous US) are generally in line with the observed averages $(4.3\pm0.5 \text{ in NCP}, \text{China and } 2.6\pm0.5 \text{ in the contiguous US})$

(Fig. 3), although the model shows a moderate low bias in both countries. -The larger pH difference in the US than in China

is likely due to the low bias in TNH₃ to which the sensitivity of pH is found to be more pronounced in the US than in China (discussed in detail in Text S1). For $c(NO_3^{-})$, the model performs generally well, with high $c(NO_3^{-})$ in China and low $c(NO_3^{-})$ in the United States (Fig. S4f&Fig. S4vi, Fig. S5f& Fig. S5vi), although c(NO₄-) in both countries are biased low (NMB is 28% in China and -30% in the United States) due to the lower bias in NO₃⁻ than in HNO₃. For c(NH₄⁺), the model reproduces the low levels in NCP and the high levels in the northeastern United States, but in China, c(NH4⁺) levels are biased high (NMB=52%) because of the low bias in NH₃ (Fig. S4g&Fig. S4e, Fig.S5g&Fig.S5e). The pH values predicted by the model are reasonable justified because 90% of the cases have the absolute pH differences between observation and simulation smaller than 1.2 in China and 2.0 in the United States. Both the NMBs and NRMSEs for pH are smaller in China than in the United States (Fig. 3).

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With respect to the temporal variation, the model captures the seasonal trends of pH, $\varepsilon(NH_4^+)$, and $\varepsilon(NO_3^-)$ in both countries. all of which are with lower values in summer and higher values in winter (Fig. 4). The lower temperature in wintertime favors the partitioning toward particle -phase for semi-volatile species. Comparison of the seasonal trends of the individual aerosol components shows a better agreement in the United States US than in China. For example, the simulation in the US captures the trends of almost all the components, though it shows biased higher concentrations of SO_4^{2-} and NH_4^{\pm} 300 than the observed ones in summer (Fig. S6, b and, h); the simulation in China misses the peaks of SO_4^{2-} in winter and NH₃ in summer, and <u>has high</u> biases for HNO₃ in summer (Fig. Figure Fig. S6, a, i, and e). On the other hand, the simulation in the United States captures the trends of almost all the components though is biased low for SO42- and NH4+in summer (Fig. S6b, h). These results indicate the need for better quantification of the monthly emission trends in China which are currently 305 subject to high uncertainty. Measurement-related biases may contribute to the disparity in the temporal trends between observed and modeled concentrations. The uncertainty in monthly profiles of emission estimates may also play an important role. For example, CASTNET'sthe long durationsampling period of the CASTNET sampling system in the US could lead to a larger measurement bias in summer than in winter (Sickles and Shadwick, 2008); The large uncertainty in the current estimates of NH₃ emissions in China, especially the reported underestimation of summertime emissions as indicated by an

- 310 inversion analysis (Kong et al., 2019), may cause the absence of the summertime NH₃ peak in the simulated trend (Fig.FigureFig. S6i). Further investigation is needed to better understand the factors underpinning the disparity between observations and model simulations. In spite of the various potential uncertainties, ooverall, the spatial and temporal evaluation suggests generally good agreement between the model simulations and observations in both countries.
- In line with the pH comparison based on observational data observations (Sect. 3.1.1), the nationwide model simulations 315 show significant differences in aerosol acidity between the two countries. Almost all the areas in the United States US have aerosol pH values lower than 3-3 according to the CDF plot-(Fig. 2b). Higher pH values are found in the middle and eastern United States US, while in the western United States US except California, the pH values are lower (Fig. 3). In China, a large portion of areas (87%) have aerosol pH values above 3-3 according to the CDF-plot, which This is especially true in the eastern China with which has the largest population (Fig. 3). Aerosol pH values in the-western and southeastern China are 320 generally lower than in the east. It should be noted that due to limitations inin the scarcity of observational data, the pH

predictionsestimates in the-southern and western China are not evaluated. The nationwide annual average pH values in China and the <u>United StatesUS</u> are 2.7±0.6 and 0.8±0.8 <u>units</u>, respectively, lower than the observation-based values <u>due partly to</u> the model bias but also because most of the monitoring sites are in the areas with high pH-areas (Fig. 3) and the bias in model simulation (Fig. 4a, Fig 4b).

Given the adverse health impacts of ambient aerosols (Burnett et al., 2014; Freedman et al., 2019) and the potential linkage of aerosol acidity with aerosol toxicity through the solubility of redox-active metals (Oakes et al., 2012; Fang et al., 2015; Ye et al., 2018), we further calculate and compare the population-weighted averages of the aerosol pH in the two countries in order-to highlight the pH levels in densely populated areas. The calculation shows the population-weighted pH values of are 3.3±0.4 and 2.2±0.5 in China and the United StatesUS, respectively, both of which are higher than non-weighted averages, which means indicating that aerosols in more populous areas tend to be less acidic (Fig. 2b). This finding is further confirmed by the statistically significant positive correlation (α=0.01), within each country, between the aerosol pH and population density (China: r=0.42, p<0.0001; the United StatesUS: r=0.28, p<0.0001). Consistent with the observation-based results, the t-test for the model simulations shows a significant difference in either both the population-weighted or and non-weighted aerosol pH values between the two countries (p<-0.001).

335 **3.2 Potential c**Causes and effects of of the aerosol pH differences

3.2.1 Gaseous and aerosolparticle compound profiles between China (NCP) and the contiguous USnited States

We further investigate the factors leading to the pH difference. Although both observations and simulations are subject to uncertainty, we expect that the observational dataobservations should provide more direct and reliable evidence for this investigation, when available. It should be noted that the monitoring sites in China were clustered in NCP and, thus, may not 340 be representative of the entire whole of China. Table 1 summarizes the annual average concentrations of gaseous and aerosol particle species measured in China (NCP) and the contiguous United States US during the study period (China: 2017,; the United States US: 2011). For all the gaseous and ionic species (except HNO₃), the average concentrations in China (NCP) are statistically significantly higher than those in the contiguous United StatesUS. The total concentrations of WSI species in China (NCP) (34.4 μg·m⁻³) are on average six times the concentrations in the contiguous United States US (5.7 μg·m⁻³) and present a larger have greater variation, ranging from $0.2-240 \ \mu g \cdot m^{-3}$, compared to the a range of $0.1-31 \ \mu g \cdot m^{-3}$ in the 345 contiguous United States US. Similar to other studies in China (Yao et al., 2002; Pathak et al., 2009; Zhang et al., 2013; Liu et al., 2016) and the United States US (Guo et al., 2015; Feng et al., 2020), NH_4^+ , NO_3^- and SO_4^{2-} , contribute more than 80% of the total WSI concentrations in both countries. The mass fractions of individual WSIs, however, differ between the two countries (Fig. 5). In China (NCP), the dominant WSI was NO₃⁻ (34.6%), followed by SO₄²⁻ (26.3%) and NH₄⁺ (25.5%). In the contiguous United States US in 2011, SO_4^{2-} contributed nearly half of the total WSI concentration (49.4%), and the 350 contributions of NO₃⁻ and NH₄⁺ are comparable (NO₃⁻ 17.6%, NH₄⁺ 18.8%), though. Note that SO₄²⁻ and NO₃⁻ levels have been found to decreasinged dramatically along over the years, leading to decreases in NH_4^+ due to since there is less substrate to interact with NH_3 and form particulate ammonium species (Butler et al., 2016).

- As tTwo of the most predominant anions in aerosols, the concentrations of SO_4^{2-} and NO_3^{-} at the monitoring sites in China 355 (NCP) are present at 4-four and 15 times the concentrations observed in the contiguous United States US, respectively. In particular, tThe relative difference in NO_3^- between the the two countries is the most significant, compared with the differences in other WSI components. Hence, the difference of the nitrate to sulfate molar ratio (NO_3^{-}/SO_4^{2-}) is also significant in between the the two countries. The observational data show that the ratios at most monitoring sites in China (NCP) are larger than 1, and that only two sites have the ratios lower but close to 1 (0.81, 0.94); \oplus On the other hand, 27 out 360 of the 34 sites in the contiguous United States US found show a ratio lower than 1, ranging from 0.25–0.99, which are generally lower than in China (NCP). High NO₃⁻/SO₄²⁻nitrate to sulfate ratio-in China (NCP) could be caused by more efficient oxidation of NO_x than SO₂ in China to allow for larger amount of leading to greater nitrate formation as well as
- higher aerosol pH and availability of ammonia NH₃ which favor the formation process of particle nitrate (Guo et al., 2018b); Vasilakos et al., 2018). The varying ratios of NO_3^{-1}/SO_4^{2-} on aerosol could further affect aerosol liquid water uptake, which 365 will be discussed in Supplementary Information -(Text- S2)in Sect. 3.2.4.
- The most abundant cation in aerosols is NH_4^+ , and the concentration difference of NH_4^+ between the two countries (by a factor of 11) is more significant thancompared with the difference of other cations (by factors of 2-7). The average NH4⁺ level at the monitoring sites in China (NCP) is more than ten times the level in the contiguous United States US. In addition, $\varepsilon(NH_4^+)$ in China (NCP) (0.13–0.48) is approximately 50% lower than in the contiguous United States US (0.22-0.85), which 370 meansing that compared to the United States US, TNH_3 in China tends to be present more in the gas phase. Higher NH_4^+ and lower ϵ (NH₄⁺) levels in China indicate amount to a higher level of TNH₃₇ which plays has an important role influence on aerosol pH, partitioning of TNO3 and even particulate mass (see Supplementary Information for more discussion), discussed in Sect 3.2.3.
- NVCs such as Na⁺, Ca²⁺, Mg²⁺, and K⁺ are often minor components of particles but are important because of their ability to 375 neutralize acidic species in the atmosphere, such as sulfuric and nitric acids (Zhang et al., 2007). Neglecting NVCs would makes cause low biases of pH, driving the NH₃-NH₄⁺ equilibrium to shifting toward the particle phase because more ammonium would be is used to neutralize the aerosol acidity than that shwould otherwise be neutralized by NVCs (Guo et al., 2018a). Therefore, NVCs are included in calculating aerosol pH in this study. High NVC concentrations usually occur at the sites near emission sources. For example, high concentrations of Na⁺, mainly from sea salt (Zhang et al., 2011), occur at 380 Sites 13, 27, and 30 in the United States US, which are all coastal sites. The High concentrations of Ca^{2+} , mainly from mineral dust, are found in greater abundance at Sites 6, 11, and 23 in the contiguous United States US and at Site 5 in China (NCP), which are are in prairies impacted by sand and dust. Average NVC concentrations in China (NCP) are up to an order of magnitude higher than in the contiguous United States US, although in both countries, most of the NVCs concentrations are small compared to SO_4^{2-} , NO_3^{-} , and NH_4^{+} . The predominant NVCs in China (NCP) are Ca^{2+} (2.8%), K^+ (2.1%), and Na^+ (2.0%) and K^+ (2.1%), while and in the contiguous United States US, they are Ca²⁺ (5.9%) and Na⁺ (3.7%).
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3.2.2 Characterization of contributions to aerosol acidity by each component individual factors

We use MTSM as described in Sect. 2.4 to characterize the contribution of each component to the pH difference between the <u>United_StatesUS</u> and China. Three groups (i.e., observation, simulation non-weighted, simulation population-weighted) of the annual average concentrations in the <u>United_StatesUS</u> and China listed in Table S4 are chosen as the starting (the <u>United_StatesUS</u>) and ending (China) points-and to perform the analysis. +The results are shown in Fig. 6.

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The average concentrations based on the observational and simulated data are not completely consistent due to the representativeness of the monitoring sites and the discrepancy between the model-simulations and observations. The MTSM analyses based on the three groups, however, showed similar results, such as that. For example, all the three groups suggest the high TNH_3 in China as an important factor leading to the difference in aerosol pH between the two countries (Fig. 6).

- 395 The contribution of TNH₃ is the highest in the "observation" group due to the large difference in TNH₃ concentration. Other cations, mainly NVCs, have a relatively small effect (0.2, 0.2, and 0.3 in groups "observation", "simulation", "simulation—weighted", respectively), which is consistent with a previous study (Zheng et al., 2020). The effects of other NVCs like Na⁺, K⁺, Mg²⁺ and Ca²⁺ on the pH difference are also considerable. The NVCs in aggregate show contributions of 0.8, 0.6, 1.0 in the "observation", "simulation", and "population-weighted simulation" groups, respectively, suggesting that the difference in the "observation".
- 400 NVCs explains approximately 1 unit difference in aerosol pH between the two countries. The higher NVC contribution in the "weighted simulation" group than in the "non-weighted" group may be explained by the interactions between NVCs and anthropogenic emissions which have been found to result in a larger control of NVCs over aerosol pH (Guo et al., 2018a; Wong et al., 2020).

Unlike TNH₃ and NVCs which lead to higher pH values in China than in the <u>United StatesUS</u>, SO_4^{2-} contributes <u>in the</u> oppositel<u>y direction</u> to the pH difference-between the two countries. High SO_4^{2-} concentrations decrease aerosol pH in China by 0.6–1.3 units-(varying by group), compared to the <u>United StatesUS</u>, although this effect is fully offset by TNH₃-and NVCs.

Compared to other species, the concentrations of TNO₃ are the most different between the two countries (by a factor of 15), but MTSM shows that the contribution of TNO₃ onto the pH difference is relatively small (0.1, 0.1, and 0.2 in the

410 observation, non-weighted, and population-weighted groups). This result is further confirmed by a sensitivity test of TNO₃ (Fig. S10) which shows that the change in pH from changing only TNO₃ is subtle small in the two-both countries with thea change in TNO₃ only. More detailed analyses and discussions on the effects of TNH₃, TNO₃, and SO₄ on aerosol pH can be found in Supplementary Information.

Studies have identified an important role of temperature in driving aerosol pH (Battaglia et al., 2017; Tao and Murphy, 2019;

Jia et al., 2020). Our MTSM analysis showeds that temperature accountsed for 0.07–0.39 unit of pH difference between China and the US, which varies by group (observation: 0.31; non weighted: 0.39; population weighted: 0.07) (Fig. 6). Such relatively small contributions of temperature, compared to those of TNH₃ and TSO₄, are mainly because of the small difference in temperature between these two countries (theywhich are located on at similar latitudes). The difference in the annual average temperature between China and the United States US is 1.4 K, -5.0 K, and 2.6 K in the observation, non-

- 420 weighted, and population-weighted groups, respectively (Table S43).
 - 3.2.3 Two pathways pertainingleading to the aerosol acidity difference
 - As aerosol pH is calculated as $[log_{10}(LWC) log_{10}(H_{air}^+) 3]$ -(both LWC and H_{air}^+ are in the unit of $\mu g \cdot m^{-3}$), all mechanisms through which factors affecting aerosol pH are associated with eithermust be through the modification of LWC, or that of H_{air}^+ , or both (LWC and H_{air}^+ are expressed as mass per unit volume of air, $\mu g m^{-3}$).- We use MTSM to
- 425 quantitatively separate the contributions of individual factors to the China-US pH difference into the LWC-modifying pathway and the H_{air}^+ -modifying pathway (Fig. 6). To achieve this, we use MTSM to quantify the contributions of individual factors to the differences in log₁₀(LWC) and [-log₁₀(H_{air}^+)-3], respectively, between the two countries, with the same approach as we did for pH (LWC and H_{air}^+ are two output variables directly predicted by ISORROPIA). The results show that both the changes in LWC and H_{air}^+ lead to increases in aerosol pH when changing the conditions change from those in
- 430 <u>the US to China.</u>

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Due primarily to the positive response of Given that LWC increases withto aerosol mass concentration (Song et al., 2019), higher component concentrations in China than in the US increase LWC and, thus, increase aerosol pH (Fig. 6b). Through such the LWC-modifying pathway, changes in SO₄, TNH₃, and TNO₃ lead to compatible increases in pH (0.15–0.3) (Fig. 6b), while, which are consistent in all three groups. Compared to other groups, -tthe observation group represents a larger

435 pH increase due to Cl and a higher pH decrease induced by Cl and due to RH, respectively, than in the other groups (Fig. 6b), mainly because of the larger differences in Cl concentrations and RH indicated byfor this group than byfor the other groups (Table S4).

Through the H_{air}^+ -modifying pathway, the effects of individual factors on pH changes are divided vary (Fig. 6c). Increases in acidic components (SO₄ and TNO₃) in China-increase H_{air}^+ -concentrations and thus decrease aerosol pH (Fig. 6c). Increases

440 <u>in TNH₃, TCl, and NVCs, on the other hand, decrease H_{air}⁺-concentrations and increase aerosol pH (Fig. 6c). By increasing H_{air}⁺-concentrations, increased SO₄ decreases pH by 0.7–1.2 units, showing a much stronger acidic capacity than another acidic component, TNO₃, which only decreases pH by 0.17–0.27 units (Fig. 6c). Compared to the <u>United States</u>US, China is in a TNH₃-rich condition. The molar ratios of [TNH₃] / (2×[SO₄]+[TNO₃]+[TCl]) in China vs. in the <u>United States</u>US are 3 vs. 1.4, 2.0 vs. 1.0, and 2.4 vs. 1.5 in the observation, non-weighted, and population-weighted groups, respectively.</u>

445 <u>Changing the conditions from the United StatesUS to China,</u> TNH₃ plays the most important role in neutralizing the acidic components and driving the pH increase in the <u>is H_{air}^+ -modifying</u> pathway (Figure Fig. 6c).

For individual factors, the net changes in pH are a result of the combination of the two pathways. For example, **L**increased SO₄ increases LWC and H_{air}^+ simultaneously. The increase in LWC increases aerosol pH, while the increase in H_{air}^+ decreases aerosol pH. All the three groups suggest that the effect of H_{air}^+ on pH overwhelms that of LWC on pH, leading to a net decrease in pH from thean SO₄ increase (Fig. 6). Increased TNH₃ increases pH in both pathways, adding up to a larger final-increase in pH (Fig. 6). Increased TNO₃ through these two pathways, however, is associated with opposite effects on

pH which are comparable in magnitude and thus tends to offset each other (especially in the observation group) (Fig. 6). This explains the aforementioned small change in pH from the TNO₃ increases. Combining all the factors, both pathways enhance increase aerosol pH (Fig. 6b and c), resulting in the large difference in aerosol acidity between these two countries

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(Fig. 6a).

- To facilitate a follow-up sensitivity test to link the two pathways with mass concentration and chemical composition, we define the total mass concentration as the sum of the mass concentrations of all the seven input components (i.e., Na⁺, SO₄, TNH₃, TNO₃, TCl, Ca^{2+} , K⁺, and Mg²⁺), including both gas and particle phases, and the chemical composition as the 460 composition of the seven components in the aerosol (gas + particle) system. The observation group shows that the total mass concentration in China is 8.4 times that in the US, and the chemical composition in China is richer in TNH₃ than that in the US (as illustrated by the ratios of $[TNH_3] / (2 \times [SO_4] + [TNO_3] + [TC])$ mentioned above). It has been found that both LWC and H_{air}⁺ are affected by mass concentration and aerosol composition (Guo et al., 2015; Xie et al., 2019; Zheng et al., 2020). To investigate how the differences in mass concentration and composition between China and the US is associated with the 465 LWC- and H_{air}-modifying pathways and consequently the pH difference, we first increase the mass concentrations of individual input components in the US case by a constant factor of 8.4, whereby we get an intervening case representing the overall pollution level as in China but with the chemical composition feature as in the US (Table S4, sensitivity test). From the intervening case, we then shift the composition of the US case to that of China (Table S4, sensitivity test). Note that throughout this sensitivity test, meteorological conditions are held constant. The first step by increasing the mass 470 concentration yields an increase of 1.02 units in the aerosol pH which is mainly achieved through the LWC-modifying pathway (1.06 units) instead of the H_{air}^+ -modifying pathway (-0.04 unit) (Fig. S7 (a), (b), (c)). The second step by changing the chemical composition shows a further increase of 0.76 units in the aerosol pH which is mainly achieved through the H_{air}⁺-modifying pathway (0.88 units), whereas the LWC-modifying pathway (-0.11 unit) plays a minor role (Fig. S7 (d), (e),
- (f)). This sensitivity test reveals that the LWC-modifying pathway is mainly associated with the change in mass
 concentration, and the H_{air}⁺-modifying pathway is mainly associated with the change in chemical composition.
- The increased pH in the first step is surprising as we initially thought that the pH should be unchanged before and after multiplying the concentration of each component with a constant factor. Further investigation shows that, during the first step, increased aerosol concentration drives more fractions of TNO₃ and TNH₃ partitioning into particle phases— ϵ (NH₄⁺) and ϵ (NO₃⁻) increase from 0.4 and 0.6 to 0.6 and 0.98, respectively. Therefore, although the chemical composition of the
- aerosol system keeps constant, the particle composition changes. This repartitioning can be explained by the Henry's Law,
 [A_{aq}]=H_Ap_A, where [A_{aq}] is the aqueous-phase concentration of A in the unit of moles per liter water, p_A is the partial pressure of A in the gas phase, and H_A is the Henry's law coefficient (Seinfeld and Pandis, 2006). In the first step after multiplying the constant factor, assuming no repartitioning, p_A increases due to the increased concentration of A in the gas phase, while [A_{aq}] remains largely constant because of the simultaneous increases in LWC and particle-phase A concentration (c_A), given that [A_{aq}] is proportional to c_A:LWC⁻¹ (note that LWC and c_A are expressed as mass per unit volume of air, and [A_{aq}] is expressed as moles per unit volume of water). According to the Henry's Law, more gas-phase A
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will dissolve in water, and thus, the total A in the system shift toward the particle phase after the multiplication of the constant factor. Due likely to the weak acidic capacity of NO_3^- which is overwhelmed by NH_4^+ , the particle is ultimately neutralized by the increased NH₄⁺ in the first step.

- 490 We find that by increasing the concentration of every component by a constant factor, the magnitude and direction of the resulting change in pH are sensitive to the fraction of TNH_3 in the aerosol system while insensitive to the ratio of SO_4 to TNO₃. Based on an NH₄⁺-SO₄²⁻-NO₃⁻-H₂O system, we conduct a series of sensitive tests to investigate the change in aerosol pH in response to the multiplication of a constant factor of 8.4 (Fig. S8). The change in pH reduces gradually from 1.2 units to 0.8 unit when the TNH₃ mass fraction of the system decreases from 67% to 27% (Fig. S8). With further decreases in the
- 495 TNH₃ fraction, the increase in pH diminishes rapidly, becomes negative when the TNH₃ mass fraction is lower than 25%, and is -0.6 when the TNH₃ mass fraction is 17% (Fig. S8). Under a constant TNH₃ mass fraction, the change in pH remain generally constant across a wide range of the mass ratios of SO₄ to TNO₃ (from 5:1 to 1:5) (Fig. S8). In populated continental regions, mass fractions of TNH₃ are often found high (Bencs et al., 2008; Behera and Sharma, 2010; Zheng et al., 2015; Cheng et al., 2016; Guo et al., 2017b), and the pH increase rather than decrease in response to the multiplication is
- 500 thus expected.
- Such an assessment by tracking pathway- and step-specific contributions gainsprovides a better understanding of the pH difference between China and the US. We show that through the LWC-modifying pathway, the increases in aerosol components consistently lead to increases in pH, and that through the Hair⁺--modifying pathway, the effects of different components on pH vary in signdirection. If- the LWC-modifying pathway dominated the pH changes over the H_{air}⁺-505 modifying pathway, aerosol mass concentrations would be the main factor driving the aerosol acidity difference between China and the US, and one could simply attribute such athe difference in aerosol acidity to the fact that China were is more polluted than the US. In contrast, if the H_{air}⁺-modifying pathway dominated, chemical composition would be the dominant factor, and the compound profiles of precursors emissions, which affect the availability of the corresponding aerosol components in the air, would play an important role. While there has been a-debate about, between whether mass 510 concentration and or chemical composition plays a more important role in determining aerosol pH (Cheng et al., 2016; Guo et al., 2017a; Pye et al., 2020; Zheng et al., 2020), our results suggest that both are important in explaining the China-US pH difference (Fig. 6b and c). The three groups are not consistent with each other in which pathway contributes more than the other to the pH difference, but they all suggest that effects on pH associated with the two pathways are comparable in terms

of their effects on aerosol pH (Fig. 6b and c).

515 Our results, showing the importance of both mass concentration associated with LWC and chemical composition associated with H_{air}^+ and a minor role of temperature, seem in some aspects to contradict a previous study (Zheng et al., 2020) which highlighted LWC and temperature instead of chemical composition as the most important factors explaining the pH difference between China (NCP) and the United States US. We note that the difference in the conclusions is reasonable when considering the differences in the specific settingsparameters cases examined in these two studies v cases is the prervious study compared the conditions in NCP in winter with those in the southeastern United States US in summer (SE-US).

Because of the differences in latitude (north for China vs south for the United StatesUS) and season (winter for China vs summer for the United StatesUS), the difference in temperature between their scenarios (29 K) was onean order of magnitude greater than those in our study which has largergreater spatial and temporal coverages (2.6 K in the observation group, 5 K in the non-weighted group, and -1.4 K in the population-weighted group). Using MTSM, we evaluate the pH

525 <u>difference between NCP and SE-US scenarios adopted</u>considered in the previous study. The results show that temperature accounts for 1.3 units of difference in aerosol pH between their two scenarios (Fig. S9), in line with what was previously reported (1.6 units).

In addition, ISORROPIA simulations show a LWC difference of 8.2 μ g·m⁻³ between China (NCP) and the contiguous US in the group-"observation" group in our study and 340 μ g·m⁻³ between the scenarios adopted considered in the previous study.

- 530 The much larger LWC difference in the previous study than that incompared to ours is mainly driven by the differences in pollutants concentrations. For example, the SO₄ concentration is as high as 156 μg·m⁻³ in the NCP scenario in the previous study, while but only 9.2 μg·m⁻³ in ours study. Such differences in concentrations are reasonable, given that the previous study selected a severe haze event occurring in Beijing in winter 2013 as the scenario for China (NCP), while we use annual average levels over NCP in 2017 as our case for China (NCP). Note that winter 2013 was a period when air pollution
- 535 reportedly reached record high levels across the northern China (Wang et al., 2014b; Li et al., 2016). Since 2013, China has launched strict controls on air pollutant emissions, and PM_{2.5} levels have decreased significantly between 2013 and 2017 (Zhang et al., 2019). Therefore, the NCP scenario in the previous study should be more representative of short-term haze events in the pre-2013 period, while our China (NCP) case should be more representative of annual average levels in recent years.

540 <u>4 Conclusion and implications</u>

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Based on extended ground-level measurements and regional air quality model simulations, we find significant differences in aerosol pH between China and the United StatesUS. Aerosols in the United StatesUS are on average more acidic with pH generally 1–2 units lower than in China. We propose an MTSM method to identify the key factors leading to the pH difference. The MTSM analysis reveals the important role of TNH₃ in causing the pH difference and an oppositeng effect from SO₄ which partially offsets the positive effect of TNH₃ on the pH change. Other factors play relatively minor roles. Further investigation highlights two pathways, one associated with changes in LWC and the other with changes in H_{air}⁺_x respectively, linking to the pH difference. The increased mass concentration in China, compared to the United StatesUS, enhances LWC, and the change in chemical composition toward a TNH₃-rich condition reduces H_{air}⁺. Both pathways facilitate the increases in aerosol pH in China and are comparable in terms of driving the pH increase.

550 <u>Our results, showing the importance of both LWC and chemical composition and a minor role of temperature, seem in some</u> aspects to contradict a previous study (Zhang et al., 2019) which highlighted LWC and temperature instead of aerosol composition as the most important factors explaining the pH difference between China (NCP) and the United States. We note that the difference in conclusion is reasonable considering the difference in the specific settings of the study cases. The pervious study compared the conditions in NCP in winter with those in the southeastern United States in summer (SE-US).
 Because of the differences in latitude (north for China vs south for the United States) and season (winter for China vs summer for the United States), the difference in temperature between their scenarios (29 K) was one order of magnitude larger than those in our study (2.6 K in the observation group, 5 K in the non-weighted group, and -1.4 K in the population-weighted group). Using MTSM, we evaluate the pH difference between the scenarios adopted in the previous study (i.e., NCP and SE US). The results show that temperature account for 1.3 units of difference in aerosol pH between their two scenarios (see SX below), which is in line with what was reported (1.6 units)."

In addition, ISORROPIA simulations show a LWC difference of 8.2 µg·m⁻³ between China (NCP) and the contiguous United States in group "observation" in our study and 340 µg·m⁻³ between the scenarios adopted in the previous study. The much larger LWC difference in the previous study than that in ours is mainly driven by the differences in pollutants concentrations. For example, the TSO₄ concentration is as high as 156 µg·m⁻³ in the NCP scenario in the previous study, while only 9.2 µg·m⁻³ in ours. Such differences in concentrations are reasonable, given that the previous study selected a severe haze event occurring in Beijing in winter 2013 as the scenario for China (NCP), while we use annual average levels over NCP in 2017 as our case for China (NCP). Note that winter 2013 was a period when air pollution reportedly reached record high levels across the northern China (Wang et al., 2014b; Li et al., 2016). Since 2013, China has launched strict controls on air pollutant emissions, and PM_{2.5}-levels decreased significantly between 2013 and 2017 (Zhang et al., 2019). Therefore, the

NCP scenario in the previous study should be more representative of short term haze events in pre 2013 period, while our

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China (NCP) case should be more representative of annual average levels in recent years.

3.2.3 Effects of ammonium on aerosol pH

The result of MTSM indicates that the difference in TNH₃ is one of the predominant reasons causing the pH difference. In order to study the effect of TNH₃, we conduct sensitivity tests for China and the United States separately to investigate the responses of aerosol pH to changing TNH₃. We change the TNH₃ concentrations from 0.1 to 1000 µg·m⁻³ while keep all other components constant at their annual average levels based on observation data (Table 2). We also use simulation data with population as the weight to study the effects, which consider other areas in China where TNH₃-concentration is not as high as in NCP. The results are shown in Fig. 7. It is clearly illustrated that, over a large range of TNH₃ concentrations, aerosol pH increases with the increase in TNH₃ because the production process of NH₄⁺ from NH₃ consumes aqueous H⁺. However, in both countries, aerosol pH has a small decrease with the increase in TNH₃ when TNH₃ concentration is very low, this could be due to higher biases in H⁺-concentration by ISORROPIA in ammonia poor conditions (Ansari and Pandis, 1999). The local sensitivity of pH to TNH₃, expressed as the pH increase per tenfold increase in TNH₃ level, is higher in the United States (3.0 based on observational data and 1.59 based on simulation data), indicated a higher sensitivity of

aerosol pH to TNH_3 in the United States than in China. Besides, we find that the responses of pH to TNH_3 are nonlinear and anisotropic. With all others equal, pH in the United States could be closer to the level in China if the TNH_3 increases to the level in China. On the other hand, the pH in China would be lower than the United States if the TNH_3 -decreases to the United States level because of the relative higher abundances of acidic components (SO₄, TNO_3 , TCl) than basic ions (TNH_3 ; NVCs) (Fig. 7a). In both countries, the sensitivities would quickly diverge from the original values toward higher values as TNH_3 decreases, with the sensitivities in China changing at a faster pace. As TNH_3 -increases, however, the sensitivities in these two countries would gradually become constant, stabilizing at comparable levels (0.002 pH unit per TNH_3 -increase in both two countries). Results based on simulation data are similar with results based on observational data, especially the sensitivity of aerosol pH at high level of TNH_3 -(represented by similar slope). Higher pH values in China based on simulation data at low TNH_3 level (1–10 µg·m⁻³) could be caused by lower SO_4^2 - concentrations. However, lower value of aerosol pH at high level of TNH_3 (> 50 µg·m⁻³) based on simulation data even with lower SO_4^2 - concentrations indicates the limit effect of TNH_3 at this level and potential effect of other components.

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The effects of TNH₃-on the gas particle partitioning of NH₃-NH₄⁺ and HNO₃-NO₃⁻ are illustrated in Fig. 7b and 7c, showing a decreasing trend of c(NH₄⁺) and an increasing trend of c(NO₃⁻) as TNH₃-increases. In the range of observation cases the value of c(NH₄⁺) in China is smaller than in the United States, suggesting excess presence of TNH₃-compared to other aerosol components (e.g., TNO₃- and SO₄). c(NO₃⁻) increased with increased TNH₃, due to higher aerosol pH which promote TNO₃ shifting to the particle phase as well as increased NH₄⁺ promote the condensation of HNO₃- to form NH₄NO₃. Higher c(NO₃⁻) in China than in the United States with an average c(NO₃⁻) in China being close to 1 confirmed the excess presence of TNH₃. Both the lower c(NH₄⁺) and higher c(NO₃⁻) in China estimated by the sensitivity curves are consistent with observations.

The gas to particle partitioning of NH₃ produces inorganic ammonium salt of ammonium bisulfate (NH₄HSO₄) and ammonium sulfate ((NH₄)₂SO₄) first because the affinity of sulfuric acid for NH₃-is much larger than that of nitric and hydrochloric acid for NH₃, especially when TNH₃-concentration is relatively low (Behera et al., 2013). The excess TNH₃may also react with nitric acid and hydrochloric acid to form salt of NH4NO3- and NH4Cl which will dissolve in the aerosol 610 liquid water (Zhao et al., 2016). Therefore, the ratio of [NH4⁺] to different acid ions ([SO4²⁻], [NO3⁻], [Cl⁻]) can be used to indicate the relative abundance of ammonia. To further investigate the effects of TNH₃ on pH at different levels of abundance, we divide the observation data into three groups based on neutralization condition of particle phase NH4⁺. Group A contains the observations when $[NH_4^+] < 2 \times [SO_4^2]$, when available NH_4^+ cannot completely balance aerosol SO_4 . Group B consists of the data points when $2 \times [SO_4^2] < [NH_4^+] < 2 \times [SO_4^2^-] + [NO_3^-] + [Cl^-]$ when most of the aerosol TSO_4 is in the 615 form of SO_4^2 - and excess [NH₄⁺] is available to stabilize nitrate and chloride driving the gas phase to shift to the particle phase. Group C contains the data points when $[NH_4^+] > 2 \times [SO_4^{2-}] + [CI^-]$, where available NH_4^+ is enough to balance particle phase anions. We then investigate the sensitivities of pH to TNH₃ in these three groups for China and the United States separately by changing the input TNH₃ from a median variation range (i.e. 55% to 150%) in each group in the two countries, respectively, and keeping all other components (i.e., concentrations and meteorological conditions) unchanged.

620 Note that no data in the United States fall in Group C, making up only two groups in the United States (i.e., Groups A and B). The results with average values of each group are shown in Fig. 8.

The aerosol pH increases with the increases in TNH₃ in all groups, which consist with the result of the sensitivity test in Fig. 7, but the increasing rates (i.e., the sensitivities of pH to TNH₃) and the pH levels vary among different groups (Fig. 8a). In China, Group C that represents aerosol systems with largest amount of excess NH₄⁺ shows the highest pH levels and the

- 625 flattest slopes of pH with TNH₃, suggesting a relatively low sensitivity of pH to the change in TNH₃ when TNH₃ is abundant. Group A that represents aerosol systems with insufficient NH_4^+ , shows the lowest pH levels with the slopes slightly steeper than in Group C. As TNH₃ decreases to 55%, the average pH in China in Group A can be as low as 2.3, closer to the pH level in the United States, consist with the conclusion in sensitivity test using average value only (Fig. 7a). Group B can be regarded as an intermediate group between Groups A and C. But the sensitivities of pH to TNH₃-changes in group B are the
- 630 highest among the three groups when reducing TNH₃, which could be due to the rapid increase in $c(NH_4^+)$ in this group as TNH₃ decreases (Fig. 8b), that leads to a faster loss of NH₄⁺ (Zheng et al., 2019). Note that although the relative abundance of NH4⁺ in group B is smaller than in group C, the transition from group B to group C due to TNH₃ increase does not always happen. Because if TNH₃-increase in an aerosol system with $2\times[SO_4^{2-}] < [NH_4^+] < 2\times[SO_4^{2-}] + [NO_3^-] + [CI^-]$, [NH₄+] would increase, and more TNO₂ and TCl would shift into the particle phase, leading to the increase of WSI concentration. However, 635 the average WSI concentration in group B is 55.03±46.79 µg·m⁻³ in China, significantly higher than that in group C in China
- (31.60±20.29 μg·m⁻³). LWC in group B (22.90±7.38 μg·m⁻³) is also higher than that in group C (14.37±16.85 μg·m⁻³). We find that most of the cases in group B could be identified as highly polluted cases where large amount of NH₄NO₃ is formed and dissolves in the aerosol water. As a result, despite the higher abundance of NH_4^+ in group B than group A, $c(NH_4^+)$ in group B is the highest among all the groups (Fig. 8b).
- 640 Throughout the observed cases, 85% in China are in Group C (i.e., aerosol systems with excess NH4⁺), and 55% in the United States are in Group A (i.e., aerosol systems with insufficient NH4⁺). The higher sensitivity of pH to TNH3 in group A than in group C explains why the pH sensitivity to TNH₃ increases more significantly in the United States than in China as TNH₃ decreases (Fig. 7a). Overall, the positive sensitivity of pH to TNH₃ and the different dominant groups in these two countries (Group C in China, Group A in the United States) suggest that the high abundance of TNH₃ in China increases the aerosol pH and is one of the major reasons for the pH difference between the two countries.
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3.2.4 The relationship between sulfate/nitrate and aerosol pH

Besides the effect of TNH₃ on aerosol pH discussed in Sec 3.2.3, other species, especially the acidic species which mainly include SO4- and TNO3, could also affect aerosol pH because of their effects on H+air concentration as well as on LWC (Ding et al., 2019). This effect is investigated in a sensitivity test by changing TNO₃ or SO₄ concentration while keeping other inputs constant as the average levels (Fig. 9). Similar to the MSTM results as shown in Fig. 6, elevated SO₄ significantly increases acrossl pH by increasing H^+_{air} . On the other hand, elevated TNO₃-only slightly increases H^+_{air} indicating a weaker acidity than that of TSO₄, in line with the result in a previous study (Guo et al., 2017b). This is partially due to the semivolatile property of TNO₃ (Ding et al., 2019). Notably, even in China where ε (NO₃⁻) are mostly close to 1, the variation of aerosol pH with TNO₃ (roughly equals to NO₃⁻ in this case) is also subtle. Therefore, for two systems with different moles of

- 655 SO₄²⁻ and NO₃⁻ neutralized by same moles of NH₄⁺, the system with more SO₄²⁻ will likely have a lower pH. This result indicates that higher aerosol acidity is associated with higher availability of TSO₄ rather than TNO₃, which can be confirmed by observed data in Fig. 10.
- Compared to the difference in TNO₃/TSO₄, the difference in NO₃⁻/SO₄²⁻ molar ratio is more significant due to higher aerosol pH and ammonium in China promotes TNO₃-shift in particle phase as NH₄NO₃, leading to a higher NO₃-/SO₄²- molar ratio. while low pH in the United States promotes TNO₃ stay in gas phase, leading to a lower NO₃-/SO₄² ratio. Based on 660 observation data, 74.5% of the cases in China have NO₃/SO₄² molar ratio larger than one, while only 22.3% in the United States. The different NO₃/SO₄²- ratios, as a result of the pH difference as well as TNO₃/SO₄ difference in two countries, could subsequently affect other aerosol properties, such as aerosol water uptake ability, which is one of the important reasons causing haze events in China during winter time (Xie et al., 2019; Wang et al., 2020b). Although nitrate aerosol and sulfate 665 aerosol absorbs similar amounts of water per mass (Fig. S7), heavy haze events in China are usually associated with increased LWC with enhanced RH levels under nitrate dominate condition (Wang et al., 2020b). In order to study this effect, we categorize the observation data into a nitrate rich group (Group N, where $[NO_3^-]/[SO_4^2] > 3$) and a sulfate rich group (Group S, where $[NO_3^-]/[SO_4^{2-}] < 1$) and compare these two groups under different RH conditions. The ratio 3 in group N is mentioned in lab studies and is a more typical value of nitrate rich conditions in field observations (Ge et al., 1998; Xie et al., 670 2019).
- The results in Fig. 11 show that aerosol pH values in the same groups in China and the United States have similar responses to the changes in RH. In both countries, as RH increases, the pH in group N decreases, and the pH in group S increases (Fig. 11a). Both the values and the increasing rate of LWC in group N is larger than in group S, suggesting a higher water uptake ability in nitrate rich condition, which is likely due to higher aerosol mass compared with group S as shown in Fig. 11f. The nearly two times aerosol mass in group N as in group S indicates the co-condensation effect of nitrate aerosol and LWC (Guo et al., 2017a), which suggests that NO₃⁻ formed in aerosol leads to a higher LWC due to the increase in aerosol mass, while higher LWC dilutes H⁺_{nir} and increases pH, which is favorable for more HNO₃ shifting from gas phase to particle phase and thus continually increases particle NO₃⁻ concentration. This effect will reach a balance when most of the gas phase HNO₃ is in the particle phase with enough NH₄⁺, and, therefore, c(NO₃⁻) is close to 100% in group N in the two countries (Fig. 11e). Besides, water uptake by hygroscopic aerosols increases the aerosol surface area and volume, enhancing the hydrolysis of N₂O₅ across particles and forming NO₃⁻ (Tian et al., 2018; Wang et al., 2020b);

The condition in group N usually has a higher LWC and aerosol mass, due to the mutual promotion between LWC and particle nitrate. And such a condition in group N occurs more often in China than in the United States, which is probably one of the reasons leading to high particle concentrations on hazy days in China.

685 4 Discussions and implications

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Based on extended ground level measurements and regional air quality model simulations, we find significant differences in aerosol pH between China and the United States. Aerosols in the United States are on average more acidic with pH generally 1–2 units lower than in China. We use two independent methods, i.e., the MTSM method and sensitivity tests, to identify the key factors leading to the pH difference. These two methods consistently reveal the important role of TNH₃-in causing the pH difference. The MTSM method further shows a significant contribution of NVCs on the pH difference, and t<u>The</u> sensitivity tests highlight the high nitrate/sulfate ratios as one of the important responses to the pH difference, and high

- nitrate aerosol in China will further lead to higher aerosol water uptake, which may have other effects to aerosol conditions. The nitrate/sulfate ratio depends on the emission ratio of NO_x/SO₂, the availability of cations due to the dependency of c(NO₃⁻) on TNH₃-(Fig. 8c, Fig. 9c), and other factors such as the atmospheric oxidizing capacity. Further investigation into the total emissions shows that the emission molar ratios of [NO_x]/[SO₂] are close to 3:1 in both countries (2.92 In China in 2017 and 3.12 in the United States in 2011 when assuming the emission NO_x-is in the form of NO₂), indicating that the emission difference is not the major factor leading to the nitrate/sulfate ratio difference. On the other hand, the emission molar ratio of [NH₃]/([NO_x]+2×[SO₂]) in China (0.75) is 1.6 times higher than that in the United States (0.46), which is consistent with the measured high relative abundance of TNH₃ in China and confirms that high availability of cations (mainly NH₄⁺-caused by high NH₃ emission) is one of the causes for the high nitrate/sulfate ratio in China.
- Will the aerosols in China be as acidic as in the United States as emissions are further controlled without significant reductions in TNH₃? Unlikely. Although both countries have been taking actions to cut down pollutant emissions (Pinder et al., 2007; Hand et al., 2012; Zhang et al., 2019), the reduction rates of NO_x and SO₂ emissions are quite different between the two countries (Fig. 12). In the United States, the reduction rates of NO_x and SO₂ emissions (mainly from mobile and power 705 sectors) were similar during the past two decades, while the emission of NH₃ (mainly from the agricultural sector) kept relatively constant (Fig. 12, https://www.epa.gov/air emissions inventories/air pollutant emissions trends data). The data in the monitoring sites in the United States showed a decreasing SO_4^2 -concentration over the years due to the SO_2 -emission reduction, but the reduction of NO₃⁻ is not obvious compared with SO₄²⁻ (Fig. S8). Lower SO₄²⁻ concentration could lead to a higher aerosol pH in the United States, but this effect could be buffered by partitioning of TNH₃, leading to a lower aerosol 710 pH than expected (Weber et al., 2016). Overall, significant higher SO_4^2 -concentration compared with relative stable $NO_3^$ concentration still led to nitrate to sulfate ratio smaller than one. This ratio, however, reached a value higher than 1 in 2015, four years after the period of this study (2011). In China, on the other hand, SO₂-emission reduction rate has been higher than NO_{*} reduction rate especially after the year 2012 (Fig. 12), which could lead to a higher nitrate to sulfate ratio (Wang et al., 2020b). Although we don't have yearly aerosol concentration data in China, the shift from sulfate dominant aerosol to nitrate 715 dominant aerosol has been already observed and reported by previous studies (Wang et al., 2011; Xie et al., 2019). We also collected the PM₂₋₅, SO₄², NO₃⁻ and NH₄⁺ concentrations in different years from many other studies in three major cities
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(Beijing, Shanghai, Guangzhou), which illustrate the increasing trend of the NO₃/SO₄²⁻ ratio along the years (Table S4). As

emissions of SO₂ and NO₂ are being controlled in China, NH₃ is becoming relatively more abundant, which is evident from the historical emission trends (decreasing SO₂ and NO_x emissions vs. steady NH₃ emission, (Zheng et al., 2018)), neutralizing the aerosol. Hence, aerosol pH in China will unlikely decrease as further emission control of NO_x and SO₂ is implemented.

Previous studies have suggested that low aerosol pH is associated with increased toxicity because of the increased dissolubility of transition metals in aerosol LWC, which induce airway injury and inflammation through the production of reactive oxygen species in vivo (Kim et al., 2015). The lower aerosol pH in the <u>United StatesUS</u> than in China implies that aerosols in the <u>United StatesUS</u> may be more toxic than in China after being inhaled by humans. However, this implication should be interpreted with caution because there are other known pathways through which particulate matter can harm the human bodyhumans, and the mechanisms of how particulate matter affects health have are not been completely understood (Armstrong et al., 2004). More studies are needed to address assess the health outcomes associated with the disparity in aerosol pH between the two countries.

730 Author contribution

HS initiated the research project. HS ran the model. HS and BZ designed the experiments, analysed results, and wrote the <u>initial draft of the</u> manuscript. YH, SX, <u>and ZX</u> helped with data preparation. All co-authors commented on <u>and contributed</u> to the writing of the paper.

Competing interests

The authors declare that they have no conflict of interest.

Data availability

The data presented in this manuscript and the observational data in China can be obtained from the corresponding author upon request. The observational data in China can also be obtained from the Data-sharing platform by Comprehensive Observation Network for Air Pollution in Beijing-Tianjin-Hebei and iIts **s**Surrounding Areas (http://123.127.175.60:8765/siteui/index). The observational data in the United States US can be obtained from Clean Air Status and Trends Network (CASTNET) (https://www.epa.gov/castnet) and Ammonia Monitoring Network (AMoN) (http://nadp.slh.wisc.edu/amon/)

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Figure 1.: Annual average aerosol pH at each monitoring site in China and the United States based on observational data. The arithmetic mean (midline), the interquartile range (box), and the minimum-maximum range (whiskers) are shown in the box plot.



Figure 2.: The cumulative distribution function (CDF) curves of aerosol pH in China and the United States based on (a) observed particulate and gaseous composition (solid lines) and CMAQ simulations collocated with observation sites (dashed line); (b) simulated data nationwide. In panel (b), both average and population weighted CDFs are shown.



Figure 3.: Overlay of annual mean pH calculated based on simulated concentrations (colored map) and observed concentrations (colored dots) over the study domain in (a) China and (b) the United States. Number of sites (N), normalized mean bias (NMB) and normalized root-mean-square error (NRMSE) are provided in each figure.



Figure 4.÷ Monthly average values of pH, ε(NO₃⁻) and ε(NH₄⁺) based on observed and CMAQ simulated data in China (a, c, e) and 1025 in the United States (b, d, f). <u>The error bars represent the standard deviation of all the cases in each month.</u>



Figure 5.* Annual average values of water-soluble ions (WSI) concentrations profiles in China (left) and in the United States (right).



Figure 6.: Contributions of individualeach componentscompounds and meteorological factorss to (a) total difference of aerosol pH (ΔpH), (b) the aerosol pH difference through the pathway of LWC (ΔpH_{LWC}), and (c) the aerosol pH difference through the pathway of H⁺_{air} ($\Delta pH_{H_{air}}$) aerosol pH difference between China and the United States calculated by multi-variable Taylor series method (MTSM) in Sect. 2.4. For each factor, T the sum of the contributions through the two pathways yields the net contribution of this factor to the aerosol pH. The case in the United States is chosen as the starting point and China as the ending point.



Figure 7: Values of pH, c(NH₄⁺) and c(NO₃⁻) to the change of TNH₃ from 0.1 to 1000 µg·m⁻³ while keep all other components constant at their annual average levels. The shaded areas show the TNH₃ concentration ranges that covers 75% of the observed cases in the countries, the black square and the red diamond mark the average TNH₃ levels in China and the United States, respectively.



Figure 8: Average values of pH, $c(NH_4^+)$ and $c(NO_3^-)$ when changing TNH₃-concentration from 55% to 150% in different groups. Group A: China: pH=3.25±1.16, n=141; US: pH=2.41±0.72, n=651; Group B: China: pH=3.68±1.13, n=240; US: pH=2.75±0.69, n=540; Group C: China: pH=4.22±1.57, n=1312; US: n=0.



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Figure 9: Values of pH, liquid water content and H⁺_{air} to the change of TSO₄ and TNO₃ concentration in China and the United States.









Figure 11: Values of pH, LWC, H⁺_{nir}, c(NH₄⁺), c(NO₃⁻) and dissolved mass in group N and group S under different RH conditions in China and the United States. China: group N, n=410; group S, n=470; US: group N, n=72; group S, n=1119.



Figure 12:-Yearly trend of the emissions of NH₂₇ NO_{*} and SO₂ in China (left) and the United States (right). The data in China are from studies by Zheng et al. (Zheng et al., 2018)., tThe data in the United States are from Air Emissions Inventories by United States Environmental Protection Agency (https://www.epa.gov/air emissions inventories/air pollutant-emissions trends-data).

Table 1_± Summary of the one-year average values of mass concentration of water_soluble ions (WSI), gaseous and aerosol species,
aerosol pH and meteorological parameters (as average ± standard deviation) in China and the United States during their study1060periods (i.e., 2017 for China and 2011 for the United States).

	China(n=1845)	US(n=1191)
WSI_(µg·m ⁻³)	34.4±25.5	5.7±2.2
Temperature_(K)	284.8±11.7	287.4±10.0
RH (%)	45.1±17.6	$71.4{\pm}20.9$
pН	4.3±1.2	2.6±0.7
Particle phase (µg·m ⁻³)		
SO4 ²⁻	9.2±7.1	2.2±1.3
NO ₃ -	12.1±11.1	0.8±0.9
$\mathbf{NH4}^{+}$	8.9 ± 8.0	0.8±0.5
Cl [.]	2.2±2.3	0.4±0.1
Na ⁺	$0.7{\pm}1.0$	0.2±0.2
\mathbf{K}^+	0.7±0.6	0.1±0.1
Ca ²⁺	1.0±0.1	0.3±0.2
Mg^{2+}	0.2±0.1	0.1±0.1

Gaseous phase (µg·m ⁻³)		
NH ₃	18.0±12.6	1.1±1.7
HCl	1.9±3.4	-
HNO ₃	$1.0{\pm}1.1$	1.0±0.6
Total (µg·m ⁻³)		
TNH ₃	26.5±17.2	$1.9{\pm}1.8$
TCl	4.1±4.5	-
TNO ₃	13.1±11.2	$1.8{\pm}1.1$