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4 **Molecular Insights into New Particle Formation**
5 **in Barcelona, Spain**
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27 **ABSTRACT**

28 Atmospheric aerosols contribute some of the greatest uncertainties to estimates of global radiative
29 forcing, and have significant effects on human health. New particle formation (NPF) is the process
30 by which new aerosols of sub-2 nm diameter form from gas-phase precursors and contributes
31 significantly to particle numbers in the atmosphere, accounting for approximately 50% of cloud
32 condensation nuclei globally. Here, we study summertime NPF in urban Barcelona in NE Spain.
33 The rate of formation of new particles is seen to increase linearly with sulphuric acid concentration
34 in a manner similar to systems studied in chamber studies involving sulphuric acid, water and
35 dimethylamine (DMA), as well as sulphuric acid, water and the oxidation products of pinanediol.
36 The sulphuric acid dimer:monomer ratio is significantly lower than that seen in experiments
37 involving sulphuric acid and DMA in chambers, indicating that stabilization of sulphuric acid
38 clusters by bases is weaker in this dataset than in chambers, and thus another mechanism, likely
39 involving the plentiful highly oxygenated organic molecules (HOMs) is plausible. The high
40 concentrations of HOMs arise largely from both alkylbenzene and monoterpene oxidation, with the
41 former providing greater concentrations of HOMs due to significant local sources. The
42 concentration of these HOMs shows a dependence on both temperature and precursor VOC
43 concentration. New particle formation without growth past 10 nm is also observed, and on these
44 days the highly oxygenated organic compound concentration is significantly lower than on days
45 with growth, and thus high concentrations of low volatility oxygenated organics appear to be a
46 necessary condition for the growth of newly formed particles in Barcelona. These results are
47 consistent with prior observations of new particle formation in both chambers and the real
48 atmosphere, and these results are likely representative of the urban background of many European
49 Mediterranean cities.

50



51 **1. INTRODUCTION**

52 Atmospheric aerosols, defined as liquid or solid droplets suspended in a gas, affect the climate both
53 directly by scattering and absorbing radiation, and indirectly by acting as cloud condensation nuclei
54 (CCN) (Penner et al., 2011), providing great uncertainties in estimates of global radiative forcing
55 (IPCC, 2014). Further, fine ambient aerosols (defined as those with diameter below 2.5 μm) are the
56 fifth greatest global mortality risk factor, resulting in 103.1 million disability-adjusted life year loss
57 in 2015 (Cohen et al., 2017). The number concentration of the ultrafine fraction of these (aerosols
58 with diameter below 0.1 μm , referred to as ultrafine particles or UFP) pose potentially significant
59 health risks also, due to their high concentration and surface area. The more diffuse, gas-like
60 behaviour of UFP allows them to penetrate into the deep lung and enter the bloodstream (Miller et
61 al., 2017). Ultrafine particles occur in the urban environment either as primary emissions (i.e., from
62 car exhaust (Harrison et al., 2018)) or secondarily as the product of new particle formation (NPF)
63 (Agudelo-Castañeda et al., 2019; Bousiotis et al., 2019)

64

65 NPF is the formation of aerosol particles from gas-phase precursors. NPF can be considered a two-
66 step process involving initial formation of a cluster of gas molecules at the critical diameter at
67 around 1.5 nm - the diameter at which a free-energy barrier must be overcome to allow the
68 spontaneous phase transition from gas to liquid or solid (Zhang et al., 2012), and the subsequent
69 growth of this droplet to a larger aerosol particle. The first step of this process is dependent upon
70 the stability and abundance of the clustering molecules. Sulphuric acid, water, and dimethylamine
71 (DMA), for example, efficiently form particles as the strong hydrogen bonding between the acid
72 base pair produces near negligible evaporation, much lower than the evaporation rate seen for the
73 more weakly bound sulphuric acid-ammonia-water system. Nucleation of sulphuric acid, DMA and
74 water proceeds at, or near to the kinetic limit (Almeida et al., 2013; Kurtén et al., 2008). Once past
75 this 1.5 nm diameter, condensation and coagulation will drive particle growth. Both the abundance
76 of condensable gases and their vapour pressures limit condensational growth. Vapour pressures are



77 especially important for the initial growth stages, as the Kelvin effect barrier impairs condensation
78 of more volatile species, with this condition of low vapour pressures becoming less significant as
79 the diameter of the particle increases (Tröstl et al., 2016). Once sufficiently large (>50 nm), the loss
80 processes for these particles become inefficient, resulting in a significant atmospheric lifetime. It is
81 at these diameters the climate forcing effects of these particles become most pronounced.

82

83 NPF processes happen globally, across a diverse range of environments from pristine polar regions,
84 to polluted urban megacities (Kerminen et al., 2018), and represent a significant source of CCN,
85 with 10-60% of NPF events shown to produce CCN and enhancement factors to CCN count ranging
86 from 0.5 – 1.1 (Lee et al., 2019 and references within). Strong NPF events are observed across a
87 range of urban environments, despite high condensation sinks (Bousiotis et al., 2019; Yu et al.,
88 2016), and can act as a precursor to strong haze events (Guo et al., 2014). Urban NPF occurs despite
89 extremely high condensation sinks, an effect which has only been partially explained by growing
90 understanding from recent in-depth studies (Yao et al., 2018). Recent advances in instrumentation
91 allow for the measurement of particles down to the critical diameter with instruments such as the
92 particle size magnifier (PSM), and (Neutral) Air Ion Spectrometer (NAIS/AIS) (Lee et al., 2019),
93 and mass spectral techniques for measuring the abundance and composition of neutral (Jokinen et
94 al., 2012) and charged (Junninen et al., 2010) clusters. Elucidated mechanisms with these
95 techniques involve sulphuric acid and ammonia in remote environments (Jokinen et al., 2018; Yan,
96 2018), monoterpene derived highly oxygenated molecules (HOM) in remote environments (Rose et
97 al., 2018), iodic acid in coastal environments (Sipilä et al., 2016), and sulphuric acid and DMA in
98 polluted urban environments (Yao et al., 2018).

99

100 Urban Barcelona sees frequent, strong summer-time NPF events. These events are associated with
101 high insolation and relatively high ozone when considering the whole year (Brines et al., 2014,
102 2015). This has been also reported in other urban environments such as Los Angeles, Brisbane,



103 Rome and Madrid (Brines et al., 2015). However, when considering mid-summer, maximum ozone
104 episodes are typically associated with high aerosol load, but not with NPF around the regional
105 background of Barcelona (Carnerero et al., 2019; Querol et al., 2017). Ground-level observations
106 report NPF events starting typically at midday, and either occurring in urban Barcelona and the
107 surrounding regional background simultaneously, or isolated to just urban Barcelona, or just the
108 regional background (Dall'Osto et al., 2013). Vertical profiles over urban Barcelona reveal that
109 NPF occurs at higher altitudes, and starts earlier in the day, as at a given altitude these events are
110 not suppressed by early traffic peaks contributing to particle load (Minguillón et al., 2015). Here,
111 we examine molecular level evidence for formation of particles at the critical diameter from
112 sulphuric acid in Barcelona, with possible contribution from strong bases and highly oxygenated
113 organic molecules (HOMs), as well as factors influencing subsequent particle growth.

114

115 **2. METHODS**

116 **2.1 Sampling Site**

117 The Palau Reial site in Barcelona (41°23'15" N, 2°6'53.64" E) is representative of the urban
118 background of Barcelona, located at the Institute of Environmental Assessment and Water Research
119 (IDAEA-CSIC) in the north-west of the city. Sampling was performed from a container 20 m from
120 a low traffic road, and 200 m from the nearest main road (Avinguda Diagonal). Data were taken
121 from 2018/06/28 through 2018/07/18.

122

123 **2.2 Chemical Ionisation Atmospheric Pressure Interface Time of Flight Mass**

124 **Spectrometry**

125 The Aerodyne Nitrate Chemical Ionisation Atmospheric Pressure Interface Time of Flight Mass
126 Spectrometer (CI-API-ToF) was used to make measurements of neutral oxygenated organic
127 compounds, organic and inorganic acids, bases, and their molecular clusters at high time resolution
128 with high resolving power. The ionization system charges molecules by adduct formation, such as



129 in the case of organic compounds with two or more hydrogen bond donor groups (Hyttinen et al.,
130 2015), or proton transfer in the case of strong acids like sulphuric acid (Jokinen et al., 2012).
131 Hydroxyl or hydroperoxyl functionalities are both common hydrogen bond donating groups, with
132 hydroperoxyl being the more efficient hydrogen bond donor (Møller et al., 2017). This instrument
133 has been explained in great detail elsewhere (Jokinen et al., 2012; Junninen et al., 2010), but briefly,
134 the front end consists of a chemical ionisation system where a 10 L min⁻¹ sample flow is drawn in
135 through the 1 m length 1" OD stainless steel tubing opening. A secondary flow was run parallel and
136 concentric to this sample flow, rendering the reaction chamber effectively wall-less. A 3 cm³ min⁻¹
137 flow of a carrier gas (N₂) is passed over a reservoir of liquid HNO₃, entraining vapour which is
138 subsequently ionised to NO₃⁻ via an X-ray source. This flow is then guided into the sample flow.
139 The nitrate ions will then charge molecules either by clustering or proton transfer. The mixed flows
140 travelling at 10 L min⁻¹ enter the critical orifice at the front end of the instrument at 0.8 L min⁻¹ and
141 are guided through a series of differentially pumped chambers before reaching the ToF analyser. All
142 data analysis was carried out in the Tofware package in Igor Pro 7 (Tofwerk AG, Switzerland).
143 Sensitivity of 3×10⁹ cm⁻³ was assumed based upon a prior calibration (Brean et al., 2019) and
144 comparison with a calculated sulphuric acid proxy (Mikkonen et al., 2011). Due to the high
145 resolving power of the CI-APi-ToF system (mass resolving power of 3000, and mass accuracy of 20
146 ppm at 201 m/Q), multiple peaks can be fit at the same unit mass and their molecular formulae
147 assigned. Beyond 500 m/Q, peak fitting and assignment of compositions becomes problematic as
148 signal decreases, mass accuracy decreases, and the total number of possible chemical compositions
149 increases, so peaks above the C₂₀ region have not been assigned (Cubison and Jimenez, 2015),
150 however, signals past this region tended to be extremely low. As proton transfer mostly happens
151 with acids, and nearly all HOM molecules will be charged by adduct formation it is possible to infer
152 the uncharged formula; therefore, all HOMs from here onwards will be listed as their uncharged
153 form. The CI-APi-ToF inlet was placed approximately 1.5 m a.g.l. CI-APi-ToF data is only
154 available between the dates 2018/07/06 and 2018/07/17.



155 **2.2 Particle Size and Number Measurements**

156 Two Scanning Mobility Particle Sizer (SMPS) instruments measured particle size distributions at 5
157 minute time resolution, one Long Column SMPS (TSI 3080 EC, 3082 Long DMA, 3772 CPC, TSI,
158 USA) and one NanoSMPS (3082 EC, 3082 Nano DMA, 3776 CPC, TSI, USA) measuring the
159 ranges 10.9 – 478.3 nm and 4.5 – 65.3 nm respectively. A Particle Size Magnifier (A10, Airmodus,
160 FN) linked to a CPC (3775, TSI, USA) measured the sub-3 nm size fraction. The PSM was run in
161 stepping mode, operating at four different saturator flows to vary the lower size cut of particles that
162 it will grow (defined as the point of 50% efficiency, D_{50}). This was set at 4 flows giving D_{50} from
163 1.4 to 2.4 nm. The instrument switched between saturator flows each 2.5 minutes, giving a sub-2.4
164 nm size distribution every 10 minutes. Aerosol sampling inlets were placed approximately 2 m a.g.l.
165

166 **2.3 Other Measurements**

167 Mixing ratios of non-methane VOC with proton affinity greater than H_3O^+ were made using the
168 proton transfer reaction time of flight mass spectrometer (PTR-ToF-MS 8000, Ionicon Analytik
169 GmbH, Austria). A detailed description of the instrument is provided by Graus et al., (2010) The
170 sampling set up, operating conditions, and quantification procedures are similar to those described
171 in Minguillón et al. (2016). Continual monitoring of composition and mass of submicron aerosol
172 was carried out with an Aerosol Chemical Speciation Monitor (ACSM, Aerodyne, USA) (Ng et al.,
173 2011). Ozone, NO, and NO_2 were measured by conventional ultraviolet and chemiluminescence air
174 quality instrumentation. Meteorological data were supplied by the Faculty of Physics of University
175 of Barcelona, from a nearby (200 m from the measurement site) meteorological station located at
176 the roof of an 8 floor building.

177

178 **2.4 Condensation Sink and Particle Growth Rate**

179 The condensation sink (CS) represents the rate at which a vapour phase molecule will collide with
180 pre-existing particle surface, and was calculated from the size distribution data as follows:



$$CS = 4\pi D \sum_{d_p} \beta_{m,d_p} d_p N_{d_p}, \quad (1)$$

182

183 where D is the diffusion coefficient of the diffusing vapour (assumed sulphuric acid), β_m is a
184 transition regime correction (Kulmala et al., 2001), d_p is particle diameter, and N_{d_p} is the number of
185 particles at diameter d_p . The formation rate of new particles at size d_p is calculated as follows:

$$J_{d_p} = \frac{dN_{d_p}}{dt} + CoagS_{d_p} \cdot N_{d_p} + \frac{GR}{\Delta d_p} \cdot N_{d_p} \quad (2)$$

187 Where the first term on the right-hand side comprises the rate at which particles enter the size d_p ,
188 and the latter two terms represent losses from this size by coagulation and growth respectively. See
189 Kulmala et al. (2001) for more information on calculation of coagulation sinks and formation rates,
190 and Kulmala et al. (2012) for calculation of growth rates.

191

192 3. RESULTS AND DISCUSSION

193 3.1 General Conditions of NPF Events

194 Summer NPF events in the regional background around Barcelona are associated with high
195 insolation, relatively low ozone (high compared with the rest of the year), and lower particulate
196 matter load (Brines et al., 2014; Carnerero et al., 2019). Figure 1 shows an example of a day with no
197 NPF in panel (a), referred to as “non-event” here, where two traffic-associated peaks in particle
198 number are seen during rush hours. Midday traffic peaks are also seen on certain days, but these are
199 easily distinguished from nucleation processes due to the lack of a significant <10 nm mode. Figure
200 1(b) shows a nucleation day with growth to larger sizes >10 nm, termed “full-event”, showing the
201 growth through the course of the day. These fulfil all the criteria of Dal Maso et al. (2005). Figure
202 1(c) shows a day with nucleation occurring, but no growth past 10 nm. These days are referred to as
203 “burst-event” days. Here, NPF is seen to occur, but particles fail to grow past the nucleation mode.
204 Two such events were seen in this data, and both are accompanied by a distinct mode appearing
205 beforehand at ~20-40 nm. Condensation sinks were not significantly higher than on full event days,
206 so this failure of particles to grow further cannot be attributed to condensational losses.



207

208 Figure 2 contains box plots showing condensation sink, temperature and global radiation for all 3
209 NPF types. Condensation sinks during NPF periods of both types (Figure 1(b) & 1(c)) were not
210 significantly lower than in non-event periods. Global radiation and temperature were higher for full-
211 event, most significantly for temperature. Figure 3 is as Figure 2 but for sulphuric acid, ammonia
212 and amines, and HOMs as measured by CI-API-ToF. Sulphuric acid is elevated during both full-
213 event and burst-event periods. In urban Barcelona, sulphuric acid will primarily arise from
214 oxidation of SO₂ by the OH radical, with anthropogenic emissions such as shipping emissions from
215 the port areas being significant sources of SO₂ (Henschel et al., 2013). Direct traffic emissions have
216 been shown to be a significant primary sulphuric acid source (Olin et al., 2020), but our sulphuric
217 acid data show no traffic peaks. Ammonia and amines show enhancement for full-event periods, but
218 not burst-event periods. Nucleation rates (at typical tropospheric sulphuric acid concentrations) are
219 sensitive to amine concentrations in the range of a few pptv, with enhancements to amine mixing
220 ratios past this point increasing the nucleation rate marginally (Almeida et al., 2013).
221 Concentrations of DMA and other alkylamines vary from zero to a few pptv in the boundary layer.

222

223 Barcelona has been shown to contain ppbv levels of ammonia (Pandolfi et al., 2012), arising from
224 both agriculture to the north (Van Damme et al., 2018), and anthropogenic activities such as waste
225 management and traffic, with waste management being the primary ammonia source. Highest
226 ammonia mixing ratios are found in the densely populated old city centre (Reche et al., 2015).
227 Agriculture, waste management, and traffic are also all significant sources of low molecular weight
228 alkylamines, such as DMA (Ábalos et al., 1999; Cadle and Mulawa, 1980; Hutchinson et al., 1982;
229 Ge et al., 2011), and are likely the source of amines found in this dataset. Activities such as
230 composting and food industry are especially strong sources of trimethylamine (TMA) (Ge et al.,
231 2011). A similar tertiary amine, trimethylamine (TEA) has been shown to be highly inefficient at
232 forming particles with sulphuric acid, likely due to its inability to form strong hydrogen bonded



233 pairs (Glasoe et al., 2015), and NO_3^- ionization is therefore likely not sensitive to tertiary amines.
234 As the charging mechanism of amines is similar to the mechanism by which amines form clusters
235 with sulphuric acid, the sensitivities will be roughly similar to the strength of the bonding within
236 sulphuric acid clusters, therefore, although significant levels of TMA are expected, they were not
237 seen in the CI-APi-ToF spectra, nor will they participate efficiently in nucleation. The
238 quantification of amines seen in the CI-APi-ToF is highly uncertain, with sensitivities likely being
239 dependent upon the efficiency with which an amine will form a cluster with the nitrate trimer in the
240 CI-APi-ToF, so no mixing ratios have been calculated from these signals (Simon et al., 2016). All
241 ammonia and amine signals follow similar trends, indicating similar sources (R^2 ranging from 0.34
242 – 0.89 between ammonia and amines). The relative strength of these signals are shown in Figure S1,
243 with significantly higher signals attributed to ammonia compared to amines, despite a likely lower
244 sensitivity (Simon et al., 2016). HOM concentrations were greatly enhanced during full-event
245 periods, but lower during burst-event periods, implying their necessity for growth. Further, all full
246 and burst events are associated with southerly and south-westerly air masses. Despite this,
247 concentrations of iodine and DMS-derived acids such as iodic acid (HIO_3) and methanesulphonic
248 acid (MSA) are low, indicating a small influence of oceanic emissions on particle
249 nucleation/growth.

250

251 **3.2 Mechanisms of New Particle Formation**

252 The correlation between J_5 and ion signals for sulphuric acid is plotted in Figure 4. J_5 is used here in
253 place of $J_{1.5}$ due to better data coverage. A close relationship between NPF rates and sulphuric acid
254 concentrations ($R^2 = 0.49$) are consistent with observations globally (Lee et al., 2019). This
255 relationship is not dependent upon condensation sink. These NPF rates have no dependence on
256 other ions as measured by CI-APi-ToF, including HIO_3 , MSA, ammonia, amines or HOMs (R^2 for
257 all < 0.1). This is not to say that all of these molecules are not involved in the nucleation process,
258 rather that elevations or reductions to their concentrations during nucleation periods do not have



259 significant impact on NPF rates. In the example of alkylamines, their gas phase concentration may
260 decrease due to clustering with elevated sulphuric acid (and therefore they will not be detectable as
261 free amines), and losses due to enhanced photochemistry during the periods of highest NPF rate.
262 Concentrations for these ions are shown in Figure S5. The strength of the relationship between
263 sulphuric acid and nucleation rate has been quantitatively reproduced in chamber studies involving
264 the H₂SO₄-H₂O-DMA, and H₂SO₄-H₂O-BioOxOrg system, accurately reproducing tropospheric
265 observations of nucleation rates (Almeida et al., 2013; Riccobono et al., 2014). A comparison
266 between our data and results from the CLOUD chamber is presented in Figure 5; included are the
267 H₂SO₄-H₂O, H₂SO₄-NH₃-H₂O (Kirkby et al., 2011), H₂SO₄-H₂O-DMA (Almeida et al., 2013) and
268 H₂SO₄-BioOxOrg-H₂O systems (Riccobono et al., 2014) – BioOxOrg refers to the oxidation
269 products of pinanediol (C₁₀H₁₈O₂) and OH. The relationship between sulphuric acid and nucleation
270 rate is similar to the H₂SO₄-DMA-H₂SO₄ system, with extremely high nucleation rates compared to
271 the other systems. The nucleation rate is also broadly similar also to the system involving
272 BioOxOrg. No dissimilarity is seen between the data points corresponding to full or burst type
273 nucleation, indicating similar mechanisms of formation, despite lower HOM concentrations on
274 these days. No higher-order sulphuric acid clusters, sulphuric acid-base clusters, nor sulphuric acid-
275 HOM clusters were visible in the mass spectral data, likely due to these being below the limit of
276 detection of the instrument (Jokinen et al., 2012), so cluster identity cannot be directly identified.
277 Sulphuric acid trimer stabilisation is dependent upon base abundance (Ortega et al 2012), and
278 conversely, sensitivity of nitrate CI-APi-ToF to sulphuric acid-base clusters is reduced due to the
279 high base content of such clusters (Jen et al., 2016).
280
281 To further explore the relationship between sulphuric acid clusters and the rate of nucleation, the
282 sulphuric acid dimer:monomer ratio is plotted in Figure 6. Sulphuric acid dimer roughly represents
283 the strength of sulphuric acid clustering in the nitrate CI-APi-ToF, as significant fragmentation of
284 sulphuric acid-base clusters occurs upon charging, with evaporation of water and bases from the



285 system (Ortega et al., 2012, 2014). Stronger bases result in more sulphuric acid dimer making it
286 through the system, and thus this ratio is higher for systems involving bases that form stronger
287 clusters with sulphuric acid. The dashed line represents the lower limit of sulphuric acid that would
288 be seen due to ion induced clustering (IIC) in the nitrate chemical ionisation system (Zhao et al.,
289 2010). The sulphuric acid dimer:monomer ratio seen in the CLOUD H₂SO₄-DMA-H₂O system is
290 plotted, alongside our own data from Barcelona. The ratio from our own data is seen to be much
291 lower than that for the system purely involving DMA as a ternary stabilising species. Similarly, this
292 ratio is lower than for reports of H₂SO₄-DMA-H₂O nucleation in Shanghai (Yao et al., 2018), but is
293 markedly similar to reports in central rural Germany (Kürten et al., 2016). Similar to central
294 Germany, this ratio increases at lower sulphuric acid concentrations to a ratio more similar to the
295 H₂SO₄-DMA-H₂O system. A possible explanation for this is that at higher sulphuric acid
296 concentrations, the concentrations of stronger stabilising bases are insufficient to stabilise all
297 present sulphuric acid, with the higher end of the sulphuric acid concentrations seen in this data
298 roughly equivalent to 1 pptv sulphuric acid ($3 \times 10^7 \text{ cm}^{-3} = 1.2 \text{ pptv sulphuric acid}$). We also cannot
299 account for clustering due to naturally charged sulphuric acid in the atmosphere, but ion
300 concentrations in urban environments tend to be small due to efficient sink processes (Hirsikko et
301 al., 2011). Particle formation therefore plausibly operates by mechanisms similar to the H₂SO₄-
302 DMA-H₂O and H₂SO₄-BioOxOrg-H₂O systems which may occur in parallel.

303

304 **3.3 HOMs and Growth**

305 Barcelona, as a densely populated urban agglomerate, is characterised by higher NO_x than the
306 remote conditions under which HOMs have primarily been studied (Bianchi et al., 2016, 2017;
307 Schobesberger et al., 2013; Yan et al., 2016). High insolation will also result in higher HO₂[·], as well
308 as higher RO₂[·] and other radicals. NO_x, HO₂[·] act as peroxy radical terminators, reducing the
309 likelihood of autoxidation to produce high O:C molecules, and the likelihood of RO₂[·]-RO₂[·]
310 dimerization. The NO_x-RO₂[·] reaction also produces HOMs with nitrate ester functionality (Brean et



311 al., 2019; Garmash et al., 2019; Rissanen, 2018), which tend to have higher volatilities, and are less
312 efficient at participating in early-stage NPF (Ehn et al., 2014; Lehtipalo et al., 2018), likely due to
313 intramolecular H-bonding (Elm et al., 2017). This, combined with plentiful VOCs with carbon
314 number <10 results in more volatile HOMs (see Figures S2 and S3) likely classed as LVOC or
315 SVOC (Bianchi et al., 2019; Tröstl et al., 2016). HOMs of this volatility are mostly incapable of
316 producing particles in the absence of sulphuric acid by ion-induced mechanisms observed in a
317 chamber study and the remote environment (Kirkby et al., 2016; Rose et al., 2018). Peaks due to
318 certain night-time HOMs are seen, largely C₁₀ HOMs attributable to inland air masses (Querol et
319 al., 2017). Many of these night-time C₁₀ HOMs contain nitrogen functionalities, attributable to
320 either oxidation by NO₃⁻ radicals or NO_x chemistry and were not associated with night-time NPF.
321 As shown in Figure 3, an elevated HOM concentration appears to be a necessary condition for
322 particle growth past 10 nm during NPF events. These days are associated with elevated
323 temperatures, solar radiation, higher ozone, and lower NO:NO₂ ratio. HIO₃ is also significantly
324 higher on burst-event days, likely a function of air mass trajectory, as HIO₃ would not inhibit
325 particle growth, condensational growth being a reversible, step-wise kinetic process. A recent study
326 in a remote environment reports growth rates matching condensation rates without accounting for
327 aqueous phase chemistry (Mohr et al., 2019). HOM yields are highly dependent upon temperature
328 (Quéléver et al., 2019; Stolzenburg et al., 2018). Lower temperatures result in slower H-
329 abstractions, which will result in the likelihood of an RO₂⁻ to undergo a different reaction pathway,
330 such as termination with HO₂⁻ to increase (Praske et al., 2018). This is particularly important if there
331 is a large energy barrier for the first or second H-abstraction taking place, as this will determine the
332 number of hydrogen bond donating groups, and therefore whether the NO₃⁻ CI-APi-ToF is sensitive
333 to a molecule or not. Figure 7(a) shows temperature plotted against the signal of HOMs. The
334 precursors for these HOMs are presumed to be largely isoprene, alkylbenzenes and monoterpenes.
335 The mean peak intensities assigned to alkylbenzene derived HOMs are approximately a factor of
336 two higher than those assigned to isoprene and monoterpene oxidation across this entire campaign.



337 In this data these VOCs are, with the exception of isoprene, not largely temperature dependent, with
338 many of these HOMs forming under negligible or zero insolation, and therefore very low OH
339 concentrations. These nighttime HOMs will not be derived from the oxidation of aromatics,
340 however, as rates of oxidation of alkylbenzenes by O₃ and NO₃ are negligible (Molteni et al, 2018).
341 These nighttime HOMs will therefore mostly be derived from biogenic emissions which undergo
342 more rapid nocturnal oxidation, and are likely transported inland by the land breeze during night
343 (Millán, 2014; Querol et al., 2017).

344

345 Operating under the assumption that C₅, C₇, and C₉ HOMs primarily arise from isoprene, toluene
346 and C₃-alkylbenzene oxidation respectively (Molteni et al., 2018; Wang et al., 2017), HOM signals
347 plotted against parent VOC concentration indicate their dependence upon that VOC. Here, a C₇
348 HOM is thought to follow the formula C₇H₈₋₁₂O₅₋₁₀N₀₋₂. This has been done in Figure 7(b). HOM
349 concentration is broadly dependent on VOC concentration. This is most significant for isoprene.
350 Fragmented monoterpene oxidation products will also contribute to the total number of C₉ HOMs,
351 and similarly, other VOCs can fragment upon oxidation. However, these results indicate that the
352 limiting factor in HOM production is temperature, and to a lesser degree VOC concentration.

353

354 Figure 8 shows three mass-defect plots for a non-event period, full-event period, and burst-event
355 period (event days are the same days as shown in Figure 1). Here, oxygenated volatile organic
356 compounds (OVOC) are defined as species visible in the nitrate CI-API-ToF that do not classify as
357 HOM. Here, the detailed criteria provided by Bianchi et al. (2019) cannot be applied to define
358 HOM, as knowledge as to whether a molecule is a result of autoxidation requires sound knowledge
359 of the structure of the precursors present. The criteria of both containing 6 oxygen atoms and 5
360 carbon atoms or greater, and having an O:C ratio >0.6 is applied, as these molecules will all
361 plausibly fulfil the updated criteria of “HOM”. The particular non-event day included in Figure 8
362 was characterised by lower solar radiation and temperatures than average, so lower signals for



363 oxygenated species are seen due to weaker photochemistry (i.e., OH concentration), and slower
364 autoxidation due to slower H-shift reactions (Frege et al., 2018; Quéléver et al., 2019; Stolzenburg
365 et al., 2018). The full-event day sees enhancements to smaller OVOCs and HOMs compared to the
366 non-event day, especially around 150-200 m/Q. Some of the largest peaks in the mass spectra
367 correspond to formulae seen arising from the OH oxidation of alkylbenzenes (Molteni et al., 2018;
368 Wang et al., 2017). Larger HOMs see a less significant enhancement to smaller alkylbenzene
369 derived HOMs. Most significant is the presence of many larger, unidentified HOMs >500 m/Q.
370 During full-event periods, these peaks are both more numerous, and larger. This area of the mass
371 spectrum will compromise the largest HOMs, from oxidation of large VOCs or from the $\text{RO}_2 +$
372 RO_2 dimerization reaction of two smaller RO_2 , these will likely have extremely low vapour
373 pressures and be able to drive early stage particle growth (Mohr et al., 2019; Tröstl et al., 2016).
374 The burst-event day has significantly lower concentrations of OVOCs and HOMs, and to a lesser
375 degree, their nitrogen containing counterparts (N-OVOCs and N-HOMs), with significantly fewer
376 compounds >500 m/Q. The most significant difference between the HOMs on both types of event
377 days is the C_9 and C_{10} HOMs, consistent with lower concentrations of monoterpenes on burst-event
378 days. The sulphur containing acids all have similar peak areas to the full-event day. These
379 elevations to condensable HOMs on particle formation days with growth consistent with particle
380 composition data as measured by ACSM, also showing a significant increase to organic mass
381 concentration in the late evening (around when new particles reach sizes measurable by ACSM) and
382 night on full-event days. This was not seen on days without nucleation, nor on days with nucleation
383 but no growth (see Figure S4).

384

385 **4. CONCLUSIONS**

386 We show new particle formation rates in Barcelona are linearly dependent upon the sulphuric acid
387 concentrations, and this mechanism plausibly proceeds by the formation of clusters involving
388 sulphuric acid and highly oxygenated organic molecules, with likely involvement of bases. This



389 multiplicity of mechanisms has been shown to occur in chamber studies but has not been observed
390 in the real atmosphere previously. The rate of nucleation relative to sulphuric acid concentration is
391 similar to observations of H₂SO₄-DMA-H₂O and H₂SO₄-BioOxOrg-H₂O nucleation. Nucleation
392 rates seem independent of HOM concentrations, implying that these mechanisms may occur in
393 competition with one another. The HOMs present in this study occur mostly from oxidation of
394 alkylbenzenes and monoterpenes, with a strong dependence of their concentration on both
395 temperature and VOC concentration. Concentrations of species associated with coastal and oceanic
396 sources such as MSA and HIO₃ were low, even though NPF was consistently associated with
397 southerly and south-westerly air masses, indicating that anthropogenic emissions of SO₂, (largely
398 arising from shipping emissions in Barcelona) and aromatic organic compounds are more important
399 precursors for compounds which initiate new particle formation and growth than oceanic emissions.
400 High HOM signals are seen to be a necessary condition for new particle growth past 10 nm, and it is
401 evident that HOM concentrations are dependent upon temperature and VOC precursor
402 concentration. Specifically large HOMs of extremely low volatility >500 m/Q are absent in the
403 mass spectra on these days without particle formation, as well as particle formation without growth.

404

405 These results are consistent with extensive chamber and flow tube studies on particle formation
406 from sulphuric acid, amines and HOMs, and further, nucleation rates relative to sulphuric acid are
407 similar to many tropospheric observations. Barcelona is representative of many Mediterranean
408 urban environments, with moderate pollution, influence of shipping emissions, and high insolation,
409 and thus the present study reveals the complexity of NPF mechanisms in these environments.

410

411 **DATA AVAILABILITY**

412 Data supporting this publication are openly available from the UBIRA eData repository at

413 <https://doi.org/10.25500/edata.bham.00000434>

414



415 **AUTHOR CONTRIBUTIONS**

416 RMH and XQ conceived the study, JB and DCSB carried out the APi-TOF and related
417 measurements with assistance from AA and MCM. The VOC measurements were proposed by NM
418 and collected by BT-R. JB wrote the first draft of the manuscript which was enhanced by
419 contributions from the co-authors.

420

421 **COMPETING INTERESTS**

422 The authors have no conflict of interests.

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430



431

432 **FIGURE LEGENDS:**

433

434 **Figure 1:** Example SMPS contour plots of (a) non-event day, (b) full-event day and (c) burst-
435 event day.

436

437 **Figure 2:** Box plots for days of non-event, full-event and burst-event, showing (a) condensation
438 sink, (b) temperature, and (c) global radiation.

439

440 **Figure 3:** Box plots for days of non-event, full-event and burst-event, showing (a) sulphuric acid,
441 (b) ammonia, C₂ and C₄ amines, as clustered with the nitrate dimer and trimer, and (c)
442 summed HOM concentration from C₅+. Units for ammonia + amines are normalised
443 counts, as no calibration was performed.

444

445 **Figure 4:** Formation rate (J₅) plotted against sulphuric acid monomer concentration, coloured by
446 condensation sink. Circles represent burst-events, squares represent full events.

447

448 **Figure 5:** Formation rate (J_{1.5}) plotted against sulphuric acid monomer concentration for data
449 collected from Barcelona. Tan circles represent burst-events, purple squares represent
450 full events. as well as that for the H₂SO₄-H₂O (blue inverted triangles), H₂SO₄-NH₃-
451 H₂O (yellow inverted triangles), H₂SO₄-DMA-H₂O (pink triangles), and H₂SO₄-
452 BioOxOrg-H₂O (brown diamonds) systems from the CLOUD chamber (Almeida et al.,
453 2013; Kirkby et al., 2011; Riccobono et al., 2014).

454

455 **Figure 6:** Sulphuric acid dimer concentration plotted against monomer concentration, showing
456 burst-event periods (tan circles), full event periods (purple squares), non-event periods
457 (green inverted triangles), and the ratio of sulphuric acid dimer:monomer in the
458 CLOUD chamber for the H₂SO₄-H₂O-DMA system (pink triangles) (Almeida et al.,
459 2013). Dashed line represents the lower limit of dimer concentration produced by ion
460 induced clustering in the chemical ionization unit (Zhao et al., 2010).

461

462 **Figure 7:** Influencing factors on VOC concentration, showing (A) temperature plotted against C₅-
463 ₁₀ HOM signal, coloured by global radiation, and (B) VOC concentration plotted against
464 HOM signal. These are segregated by carbon number/VOC, i.e. C₇ HOMs plotted
465 against toluene, under the assumption that toluene oxidation is the main producer of C₇
466 HOMs. Ellipses show 95% confidence on a multivariate t-distribution.

467

468 **Figure 8:** Mass defect plots for (a) non-event, (b) full-event, and (c) burst-event periods, data
469 taken from 10:00 – 15:00 on the days 12/07/2018, 16/07/2018 and 15/07/2018
470 respectively. Size corresponds to mass spectral peak area. Ions are coloured according
471 to identified chemical composition. *Blue* points correspond to HOMs containing all
472 organic species with ≥5 carbon atoms and ≥6 oxygen atoms, and an O:C ratio of >0.6.
473 *Purple* points correspond to the same but for species containing 1-2 nitrogen atoms.
474 Species not meeting this HOM criteria were classed generally as OVOCs which are
475 coloured *brown*, with the nitrogen containing OVOCs coloured *orange*. Sulphur acids
476 (*red*) include ions HSO₄⁻, CH₃SO₃⁻ and SO₅⁻, as well as the sulphuric acid dimer. Iodine
477 acids (*green*) contains both IO₃⁻ and I⁻ (the latter presumably deprotonated hydrogen
478 iodide). Unidentified points are left uncoloured.

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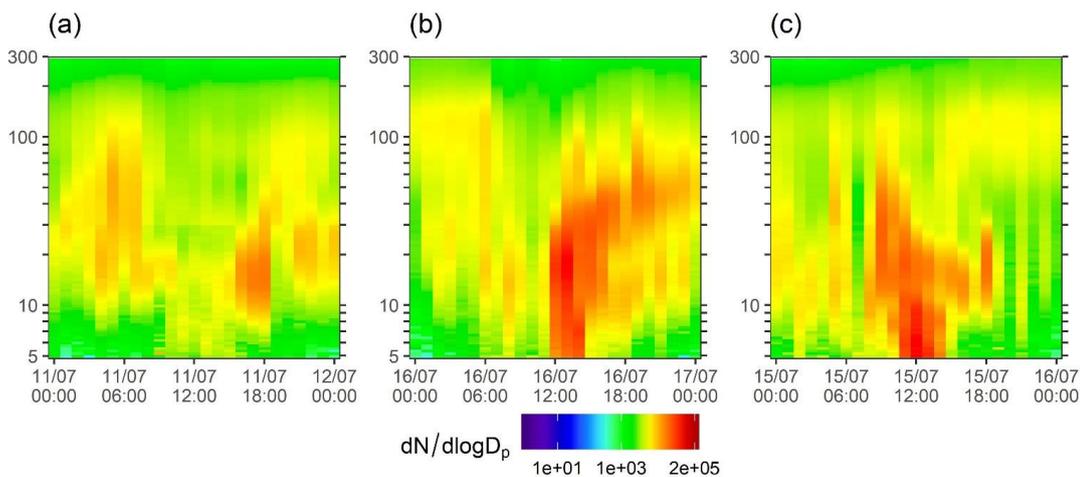
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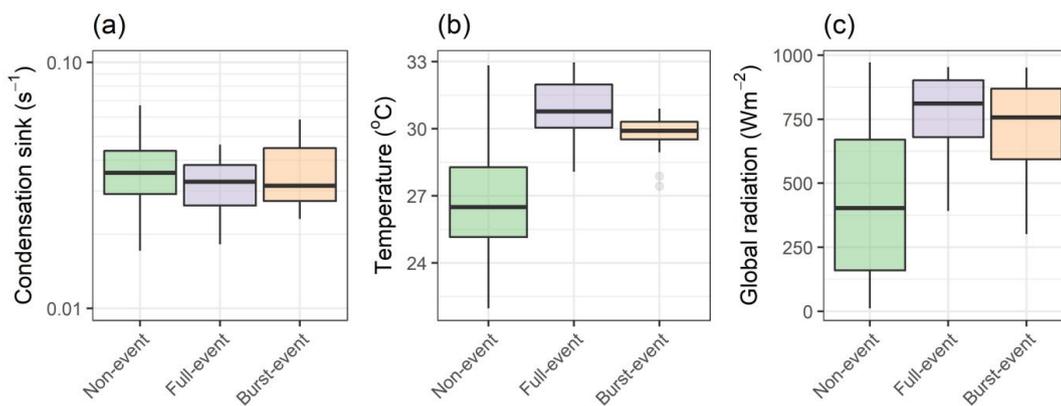


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907 **Figure 1:** Example SMPS contour plots of (a) non-event day, (b) full-event day and (c) burst-event
908 day.

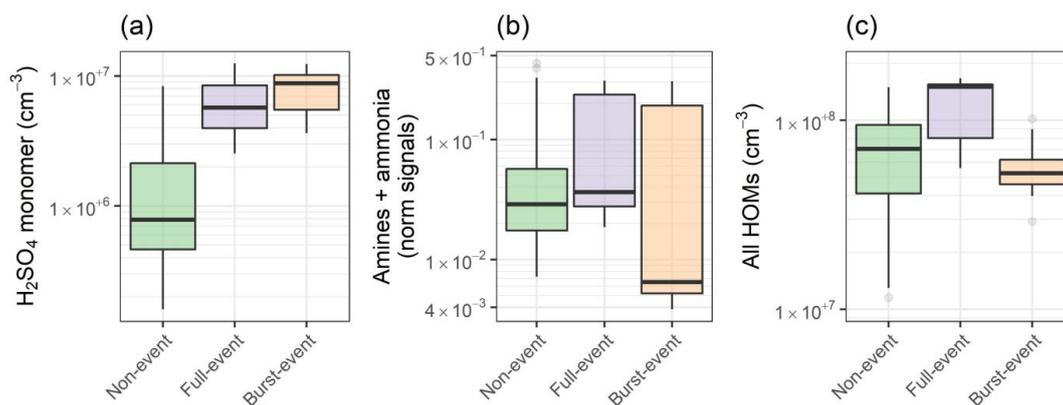
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911 **Figure 2:** Box plots for days of non-event, full-event and burst-event, showing (a) condensation sink,
912 (b) temperature, and (c) global radiation.

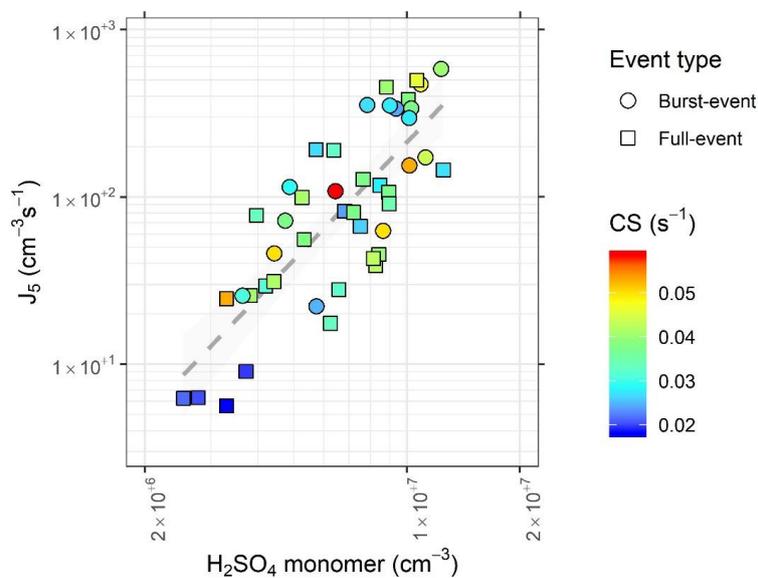
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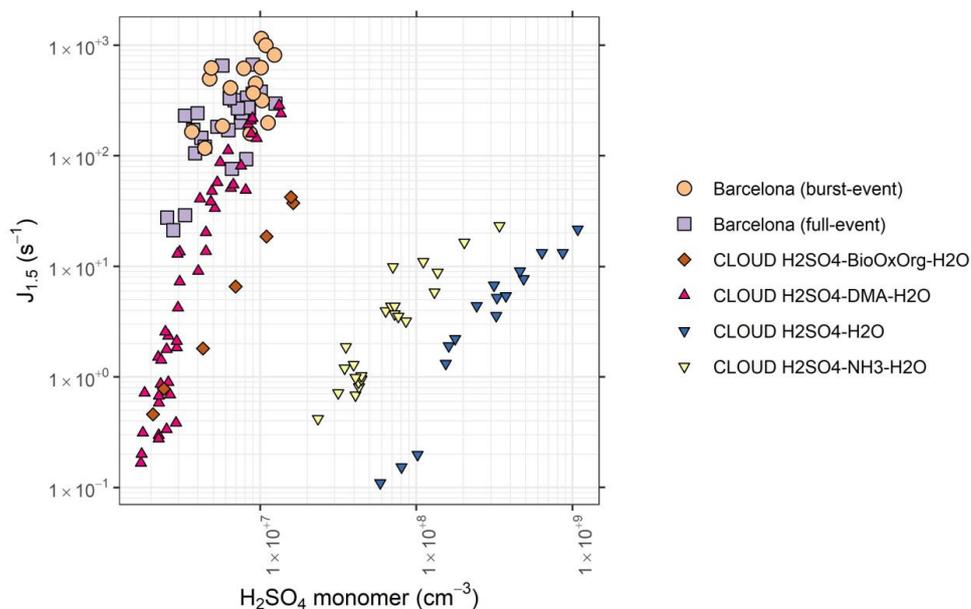
915 **Figure 3:** Box plots for days of non-event, full-event and burst-event, showing (a) sulphuric acid, (b)
916 ammonia, C₂ and C₄ amines, as clustered with the nitrate dimer and trimer, and (c) summed HOM
917 concentration from C₅+. Units for ammonia + amines are normalised counts, as no calibration was
918 performed.

919



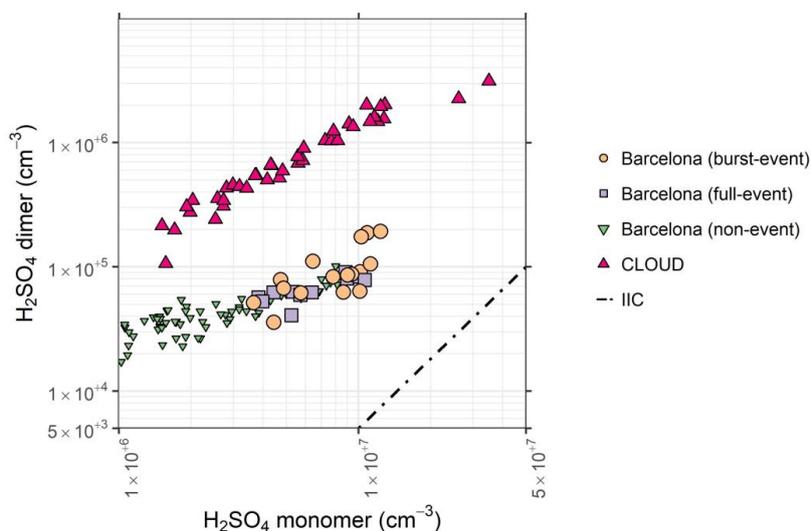
920

921 **Figure 4:** Formation rate (J_5) plotted against sulphuric acid monomer concentration, coloured by
922 condensation sink. Circles represent burst-events, squares represent full events.



923

924 **Figure 5:** Formation rate ($J_{1.5}$) plotted against sulphuric acid monomer concentration for data
925 collected from Barcelona. Tan circles represent burst-events, purple squares represent full events, as
926 well as that for the $\text{H}_2\text{SO}_4\text{-H}_2\text{O}$ (blue inverted triangles), $\text{H}_2\text{SO}_4\text{-NH}_3\text{-H}_2\text{O}$ (yellow inverted triangles),
927 $\text{H}_2\text{SO}_4\text{-DMA-H}_2\text{O}$ (pink triangles), and $\text{H}_2\text{SO}_4\text{-BioOxOrg-H}_2\text{O}$ (brown diamonds) systems from the
928 CLOUD chamber (Almeida et al., 2013; Kirkby et al., 2011; Riccobono et al., 2014).

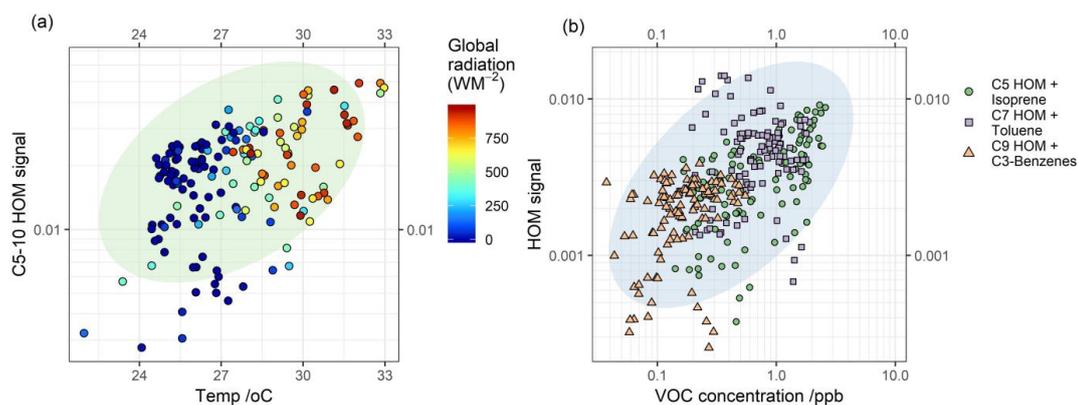


929

930 **Figure 6:** Sulphuric acid dimer concentration plotted against monomer concentration, showing burst-
931 event periods (tan circles), full event periods (purple squares), non-event periods (green inverted
932 triangles), and the ratio of sulphuric acid dimer:monomer in the CLOUD chamber for the $\text{H}_2\text{SO}_4\text{-}$
933 $\text{H}_2\text{O-DMA}$ system (pink triangles) (Almeida et al., 2013). Dashed line represents the lower limit of



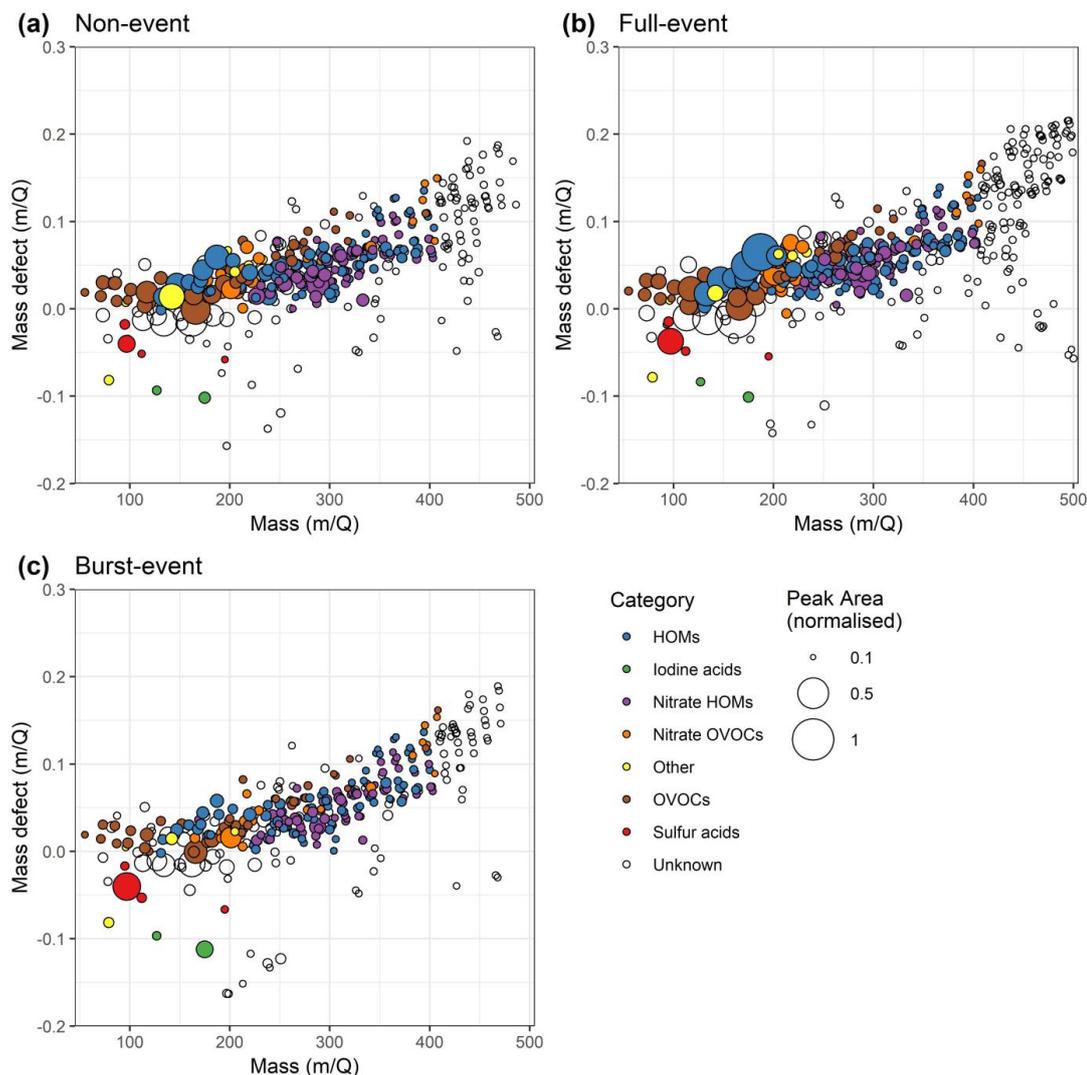
934 dimer concentration produced by ion induced clustering in the chemical ionization unit (Zhao et al.,
935 2010).



936

937 **Figure 7:** Influencing factors on VOC concentration, showing (A) temperature plotted against C₅₋₁₀
938 HOM signal, coloured by global radiation, and (B) VOC concentration plotted against HOM signal.
939 These are segregated by carbon number/VOC, i.e. C₇ HOMs plotted against toluene, under the
940 assumption that toluene oxidation is the main producer of C₇ HOMs. Ellipses show 95% confidence
941 on a multivariate t-distribution.

942



943

944 **Figure 8:** Mass defect plots for (a) non-event, (b) full-event, and (c) burst-event periods, data taken
945 from 10:00 – 15:00 on the days 12/07/2018, 16/07/2018 and 15/07/2018 respectively. Size
946 corresponds to mass spectral peak area. Ions are coloured according to identified chemical
947 composition. *Blue* points correspond to HOMs containing all organic species with ≥ 5 carbon atoms
948 and ≥ 6 oxygen atoms, and an O:C ratio of >0.6 . *Purple* points correspond to the same but for species
949 containing 1–2 nitrogen atoms. Species not meeting this HOM criteria were classed generally as
950 OVOCs which are coloured *brown*, with the nitrogen containing OVOCs coloured *orange*. Sulphur
951 acids (*red*) include ions HSO_4^- , CH_3SO_3^- and SO_5^- , as well as the sulphuric acid dimer. Iodine acids
952 (*green*) contains both IO_3^- and I^- (the latter presumably deprotonated hydrogen iodide). Unidentified
953 points are left uncoloured.

954