

RESPONSE TO REVIEWERS

We thank all reviewers for their insightful comments on our work and are pleased to respond as below:

Note: Referee comments in this text will appear in blue, responses in black, and where new text from the manuscript is quoted, it will appear in red, and **bold** where text has been inserted into a larger section

Referee 1

Major comment 1: Values for condensation sink (CS), J1.5, J5, growth rates (GR). There is a mistake in CS formula (line 181). The coefficient 4 should be 2, as in Kulmala et al. (2001).

J1.5 and J5 are very high, but in line (within some uncertainty) with calculated kinetic limit sulfuric acid nucleation (Kurten et al. 2018). However, during the observed NPF, the temperature is close to 30 degC. At this high temperature, I would expect lower formation rates. Please, check the values and provide a comment on this.

It would be beneficial, if more information was included how CoagS and GR were calculated (shortly from which instrument and with which method GR was calculated). Would it be beneficial to include GR values to results as important physical parameter of NPF?

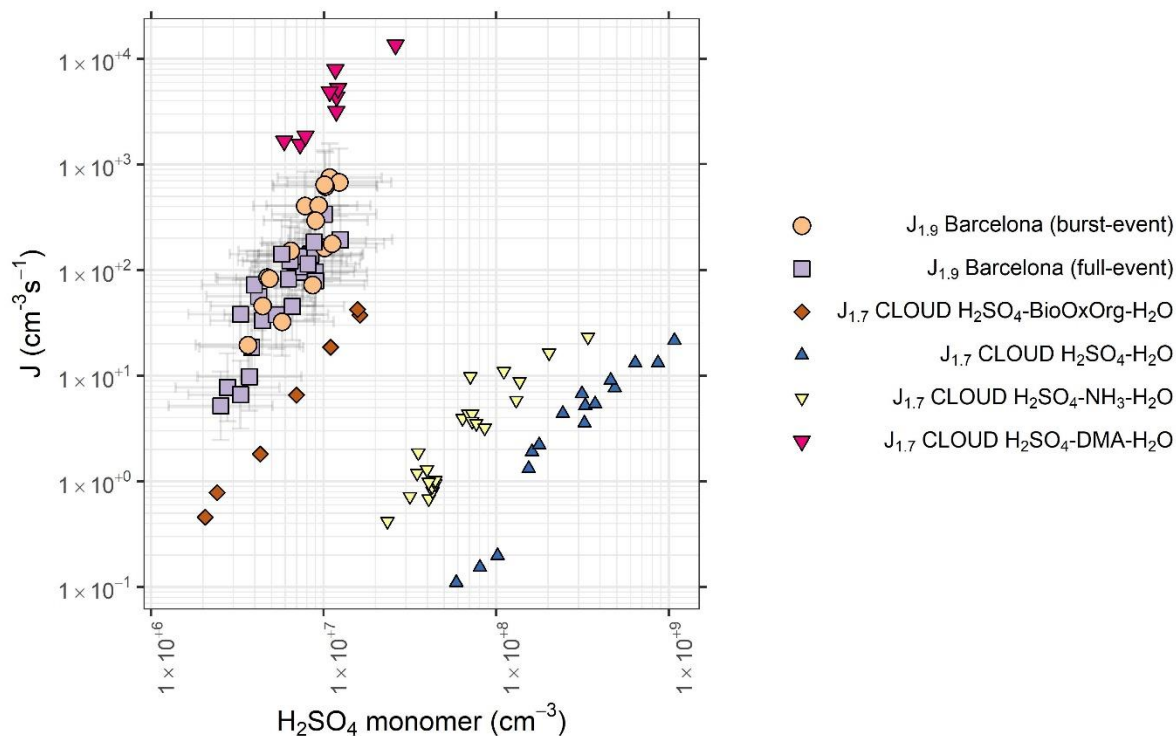
Response 1: Equation 14 in Kulmala et al., 2001 states $4\pi D$, this is consistent across more recent publications also (Kulmala et al., 2012). We have included a reference to the latter just before this equation also.

We agree our formation rates are very high – these calculations have been double checked for clarity. We note, however, that such high formation rates at high temperatures have been observed previously, examples being from Mexico City ((Iida et al., 2008; Kuang et al., 2008), temperatures of 19.2 and 24.1 °C are quoted in the former, high formation rates relative to sulphuric acid are quoted in the latter. Similarly, Kürten et al, (2016) report high formation rates on days with peak temperatures just below 25 °C. Measurements in Beijing show a rather sharp temperature dependence, but particle formation rates are shown to reach $10^2 \text{ cm}^{-3} \text{ s}^{-1}$ at temperatures similar to our own at comparable sulphuric acid concentrations (Yu et al., 2016). We include the following discussion in the text addressing this.

“Model studies of sulphuric acid-amine nucleation show a decline in nucleation rate with temperature (Almeida et al., 2013; Olenius et al., 2017), as the evaporation rate of sulphuric acid-amine clusters will increase with temperature (Paasonen et al., 2012). Conversely, evaporation rates of such small clusters, and resultant nucleation rates tend to increase modestly with increases in relative humidity, most pronounced at lower amine concentrations (Almeida et al., 2013; Paasonen et al., 2012). Despite this, high nucleation rates at temperatures nearing 300 K have been reported previously (Kuang et al., 2008; Kürten et al., 2016), although these tend to show a temperature dependence (Yu et al., 2016).”

We also utilise updated H₂SO₄-DMA-H₂O data in our Figure 5 from Kürten et al. (2018), where we show our measurements fall short of the J values from the CLOUD chamber experiments, and thus we edit our conclusions accordingly and suggest the discrepancy arises from temperature.

“while formation rates far exceed that of $\text{H}_2\text{SO}_4\text{-BioOxOrg-H}_2\text{O}$ nucleation, they fall short of those of $\text{H}_2\text{SO}_4\text{-DMA-H}_2\text{O}$ nucleation at 278 K, as does the sulphuric acid dimer:monomer ratio, possibly explained by cluster evaporation due to high temperatures in summertime Barcelona (303 K during events), and limited pools of gas-phase amines. These results are similar to reports of nucleation rates in rural Germany (Kürten et al., 2016). As no higher-order clusters were directly measured, we cannot determine nucleation mechanisms with certainty, and an involvement of HOMs in nucleation is plausible.”



Further to this, we present an expanded methods section

“ J_5 was calculated using the data between 5-10 nm, and $J_{1.9}$ was calculated using the measurements between 1.9 – 4.5 nm. We also calculated $J_{1.9}$ from our NanoSMPS data, employing the equations of Lehtinen et al. (2007). $J_{1.9}$ from both methods showed reasonable agreement ($R^2 = 0.34$). Agreement between J_5 and $J_{1.9}$ for each method was similar ($R^2 = 0.37$ and $R^2 = 0.38$ for $J_{1.9}$ calculated from PSM data and from Lehtinen et al. (2019) respectively) $J_{1.9}$ is greater than J_5 as predicted from equation (2) by around a factor of 20. See Kulmala et al. (2001) for more information on calculation of coagulation sinks and formation rates. Growth rates between 4.5 – 20 nm were calculated according to the lognormal distribution function method (Kulmala et al., 2012), whereas those between 1.9 and 4.5 nm were calculated from PSM data using a time-delay method between PSM and NanoSMPS data.”

We also include GR values in section 3.1

“ $GR_{4.5-20}$ ranged between 2.47 – 7.31 nm h^{-1} ($4.69 \pm 2.03 \text{ nm h}^{-1}$), $GR_{1.9-4.5}$ ranged between 3.12 – 5.2 nm h^{-1} ($4.36 \pm 1.02 \text{ nm h}^{-1}$).”

Major comment 2: “Burst event”/ “full-event” discussion

The discussion what separates these event types is hard to follow for the following 2 reasons

1) In Fig. 2-6, how the data was produced? Are these values during certain part of the day or is whole day included? Are these hourly values? Please specify this in Methods or Results.

For instance, in terms of CS comparison, CS over whole day may not show any difference (lines 205-206, 209-210 and Fig. 2a). On the other hand, daily profile of CS may show that on “burst event” (Figure 1c), 100nm mode appears around 14:00, which I assume increases afternoon CS during that day.

2) The discussion about differences in meteorology is incomplete. In lines 246 and 397, it is mentioned that south and south-west air masses were responsible for NPF, while in Fig. S6, air masses on 3/5 NPF days came from North. Could change in air mass (wind direction) be responsible for no growth during “burst events”? Please clarify.

3) Please provide information on which day, what type of NPF was seen.

Response 2:

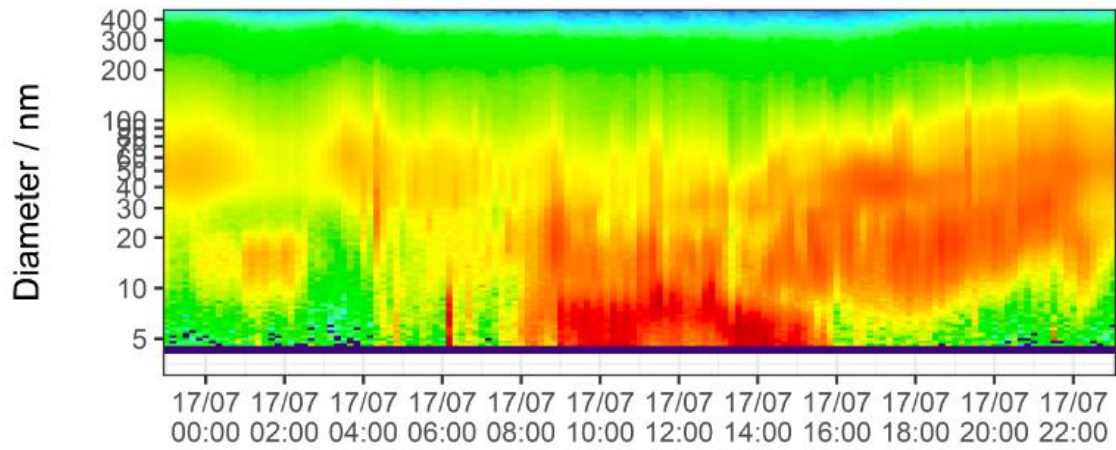
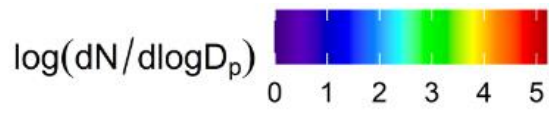
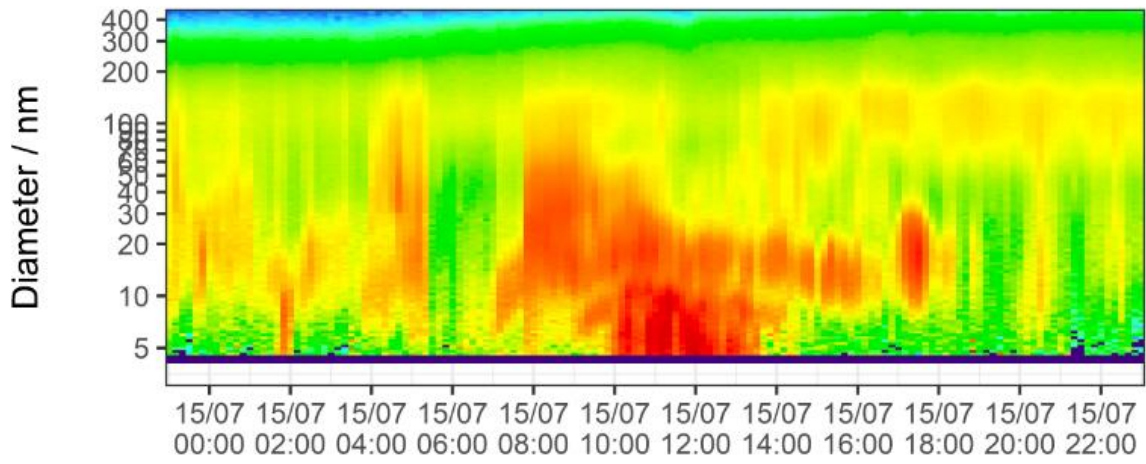
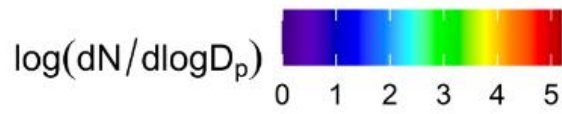
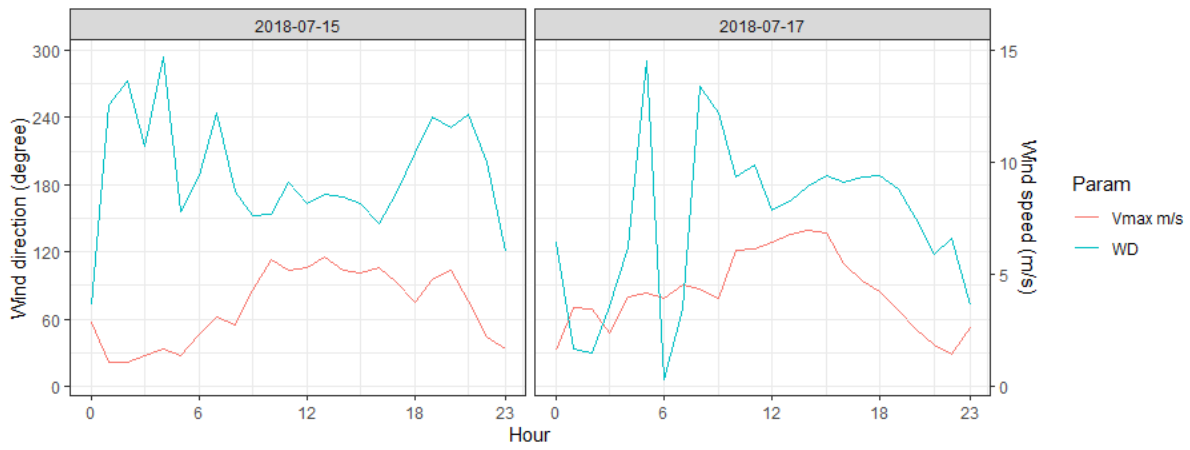
1) The periods of data that have been included have been clarified in the following sections

- Figure captions for Figures 2 + 3 include the sentence *“Event days include data across the full event day.”*
- Data coverage is a little misleading. Figures 4 – 6 have been amended for consistency, containing the phrase *“Data is for hourly averages across NPF periods, typically within the hours 08:00 – 16:00.”*

Further, we include a new Figure S2 showing the diurnal evolution of condensation sinks on non-event days, as well as both types of event days, alongside this we present an extended discussion where we include the following

“Figure 2 contains box plots showing condensation sink, temperature and global radiation for all 3 NPF types across the entire day (diurnal profiles plotted in Figure S1). Condensation sinks during NPF periods of both types (Figure 1(b) & 1(c)) were not significantly lower than in non-event periods. Condensation sinks were suppressed prior to the beginning of an event for full-events, increasing relative to non-events through the afternoon period. Of the two burst-events, one was similarly characterised by a suppression to condensation sink, whereas the other showed a sharp rise in the midday”

2) This discussion was based upon an analysis of local wind direction + speed measurements. From polar plots it was inferred that air masses arising from the south were dominant during NPF; however, we find the air mass history from HYSPLIT to be more informative to our analyses, and thus these were included. The wind direction and velocity across the burst event days is shown below, where it can be seen that air masses are stable following the appearance of the new mode, and persist across the period of multiple hours, giving us confidence that this is not due to air mass advection.



3) This information is now available in Figure S8, where air mass history is presented by day, with each class of day stated in the figure caption.

Major Comment 3: Discussion about HOM sources/formation/influence on particle growth

The presentation and discussion on HOM results is very valuable. To make the section 3.3 more clear, I suggest to separate it in two subsection: 1) HOM composition and sources and 2) How HOM change on different days in relation to NPF. The following are my main concerns regarding HOM discussion:

1) Interpretation of Figure 7. Are there really dependencies seen?

Fig. 7a to me shows correlation between T and global radiation. If night and day are separated, there is no dependency on T, or is there?

In Fig. 7b, dependency is discussed when all points are taken together. But is there a point to do that if different groups of HOMs are plotted against different VOCs? For C7/C9 I don't see any dependence.

Therefore, the conclusion that HOM has "strong dependence of their concentration on both temperature and VOC concentration" (line 394) is not supported.

In Fig. 7a HOM C5-10 are plotted, while in Fig. 7b, only C5/C7/C9 are shown. What is the reason for excluding for instance C6,C8,C10?

Please clarify this section and harmonize the text (abstract-discussion-conclusions).

2) Sudden statement in conclusion regarding HOM effect on NPF and growth:

In lines 395-399, it is stated "aromatic organic compounds are more important compounds which initiate new particle formation and growth than oceanic emissions". Which data leads to this conclusions? Especially, the part in which aromatic compounds initiate NPF. In line 377, it is actually mentioned that biggest difference between the two event types are in fact C9/C10 compounds. Please clarify this aspect.

3) Please clarify what is the effect of HOM on NPF vs growth, if such can be distinguished from the data.

Response 3:

We have separated these sections as suggested, and provide the following answers below

1) If we take just the daylight hours, or hours where insolation is $>100 \text{ W m}^{-2}$, the correlation between temperature and signal gets slightly better $R^2 = 0.29 \rightarrow R^2 = 0.30$. Taking the periods where insolation is $<100 \text{ W m}^{-2}$ then there is no correlation seen. However, the correlation between HOMs and insolation is poor ($R^2 = 0.15$), and we maintain that there exists a dependence of HOM concentration on temperature within this dataset.

We excluded C₆, C₈ and C₁₀ from this plot so as to not overcomplicate the image. In the case of C₁₀ also, we measured C₁₀ aromatics and monoterpenes at approximately a 1:3 ratio, and these will both produce oxidation products of similar chemical formulae, thus including C₁₀ in the plot comes with some uncertainties. However, we produce a new version of the plot including isoprene, benzene, toluene, C₂ aromatics and C₃ aromatics. We concur with the skepticism about our conclusions from these figures and have updated this entire section. Specifically, regarding figure 7B we write

“Operating under the assumption that C₅, C₆, C₇, C₈, and C₉ HOMs primarily arise from isoprene, benzene toluene, C₂-alkylbenzene C₃-alkylbenzene oxidation respectively (Molteni et al., 2018; Wang et al., 2017), HOM signals plotted against parent VOC concentration indicate their dependence upon that VOC. Here, a C₇ HOM is thought to follow the formula C₇H₈₋₁₂O₅₋₁₀N₀₋₂. We have plotted HOM concentrations against VOC concentrations in Figure 7(b). C₁₀ HOMs are not included in these analyses as these may primarily arise from C₁₀H₁₂₋₁₄ alkylbenzene, or monoterpene oxidation. Concentration appears mostly independent of VOC concentration, with the exception of isoprene, for which emissions are highly temperature dependent, and thus this is likely a function of the effect of temperature on HOM formation (Figure 7(a)). A lack of correlation between other VOCs and their HOMs confirms that this relationship between HOMs and temperature is not a function of enhanced source fluxes from, for example, evaporation, except in the instance of isoprene”

2) In the text we omit this statement, and provide a much revised form of these sections, as well as our conclusions re: NPF mechanisms (see response to major comment 1), this particular sentence now reads as follows.

“The burst-event day has significantly lower concentrations of OVOCs and HOMs, and to a lesser degree, their nitrogen containing counterparts (N-OVOCs and N-HOMs), with significantly fewer compounds >400 m/Q. The most significant difference between full and burst-event days is in the SVOCs, accounting for a factor of two difference in concentration. The sulphur containing acids all have similar peak areas to the full-event day”

3) The effect of HOM on NPF vs growth is discussed in the conclusions now, through the sentences

“High concentrations of OVOC and HOMs were measured by CI-APi-ToF. Of these, the SVOC arose from mostly isoprene and alkylbenzene oxidation, whereas LVOC and ELVOC arose from alkylbenzene, monoterpene and PAH oxidation, with a dependence of their concentration on temperature, but less so on VOC concentration. Concentrations of species associated with coastal and oceanic sources such as MSA and HIO₃ were low. High HOM signals are seen to be a necessary condition for new particle growth past 10 nm, with the most significant difference between days with and without particle growth being SVOC concentrations (factor of 2 difference), while modelled growth rates from condensation of these organic compounds, alongside H₂SO₄, MSA and HIO₃ were shown to match growth rates within measurement error. Thus, oxidation of traffic derived alkylbenzenes and PAHs, and to a lesser degree, isoprene and monoterpene emissions are significant for new particle growth in this environment.”

These conclusions are based upon analysis of volatility and double bond equivalency (DBE) in our measured organic molecules. We have calculated DBE per molecule and saturation mass concentration as according to 2D-VBS, explained in the methods section as follows

“The double bond equivalent (DBE) describes the degree of unsaturation of an organic molecule and is defined simply as

$$DBE = N_C - \frac{N_H}{2} - \frac{N_N}{2} + 1 \quad (3)$$

The saturation vapour pressure at 300 K is defined by the 2D-volatility basis set (2D-VBS) as follows, if all nitrogen functionality is assumed to take the form -ONO₂ (Bianchi 2019; Donahue 2011; Schervish and Donahue, 2020):

$$\text{Log}_{10}(C^*)(300\text{ K}) = (N_{C0} - N_C)b_C - N_O b_O - 2 \frac{N_O N_C}{N_C + N_O} b_{CO} - N_n b_N \quad (4)$$

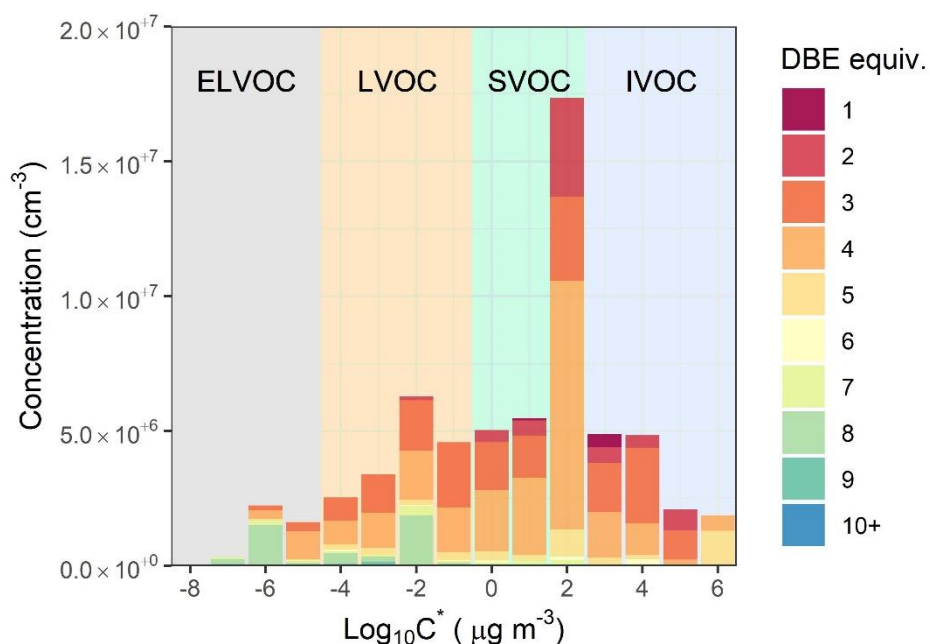
Where N_C, N_H, and N_N, are the number of carbon, hydrogen, and nitrogen atoms respectively. N_O is the number of oxygen atoms minus 3N_N to account for -ONO₂ groups, N_{C0} is 25 (the carbon number of a 1 μg m⁻³ alkane), b_C, b_O, b_{CO}, and b_N are 0.475, 0.2, 0.9 and 2.5 respectively, and represent

interaction and nonideality terms. The final term of equation (4) represents the $-ONO_2$ groups, each reducing the saturation vapour pressure by 2.5 orders of magnitude. C^* values are calculated at 300 K and not corrected, as 300 K is within 1 K of the campaign average temperature.”

We include the following sections in our discussion of HOM composition and sources

“DBE as calculated by equation 3 is equal to the number of pi bonds and rings within a molecule. Benzene, toluene, and similar aromatics have $DBE = 4$, naphthalene = 7 and monoterpenes = 3. DBE can be used as an indicator of sources when considering HOM in bulk. Saturation mass concentration as calculated by equation 4 can help describe capacity of a molecule to both condense onto newly formed particles and participate in nucleation. Figure 8 shows concentrations of HOMs and other oxygenated organic molecules binned to the nearest integer $\text{Log}_{10}(C^*)(300\text{ K})$, coloured by DBE. Mean ion signals per carbon number are shown in Figure S7. Most measured molecules fall into the SVOC class ($0.3 < C^*(300\text{ K}) < 300\ \mu\text{g m}^{-3}$) which will mostly exist in equilibrium between gas and particle phase. High SVOC concentrations arise from fingerprint molecules for isoprene oxidation under high NO_x ($\text{C}_5\text{H}_{10}\text{N}_2\text{O}_8$) (Brean et al., 2019), and oxidation of small alkylbenzenes ($\text{C}_7\text{H}_8\text{O}_5$, $\text{C}_8\text{H}_{10}\text{O}_5$). LVOC and ELVOC ($3 \cdot 10^{-5} < C^* < 0.3\ \mu\text{g m}^{-3}$ and $3 \cdot 10^{-9} < C^*(300\text{ K}) < 3 \cdot 10^{-5}\ \mu\text{g m}^{-3}$ respectively) have a greater contribution from molecules with higher DBE, i.e., $\text{C}_{10}\text{H}_{10}\text{O}_8$ arising most likely from PAH oxidation (Molteni et al., 2018), and $\text{C}_{10}\text{H}_{15}\text{O}_7\text{N}$, a common molecule arising from monoterpene oxidation in the presence of NO_x . The contribution of molecules with carbon number ≤ 9 to these LVOC is modest, and ELVOCs are entirely comprised of molecules with carbon numbers ≥ 10 and DBEs of 8 and 4. No molecules classed as ultra-low volatility organic compounds (ULVOC, $C^*(300\text{ K}) < 3 \cdot 10^{-9}\ \mu\text{g m}^{-3}$) were observed.”

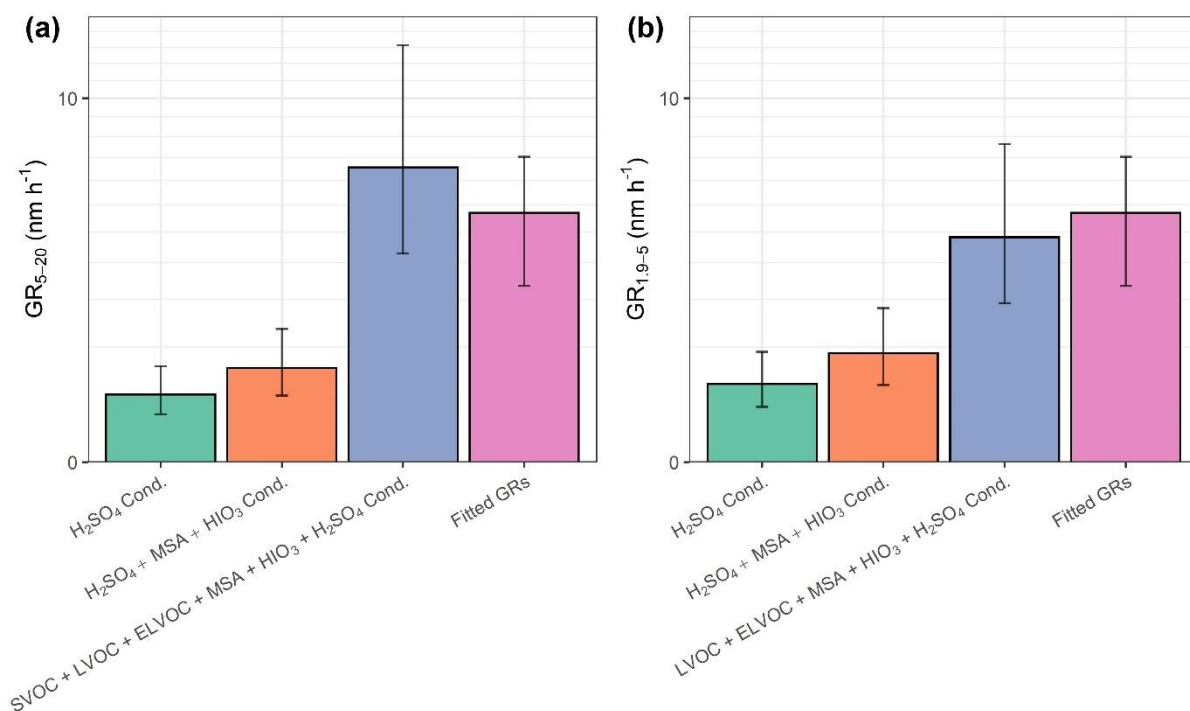
We discuss these in the context of NPF in section 3.3.2. We present our new figure below



A parametrisation of condensational growth from Nieminen et al. (2010) has been applied to these results, included in the supplement (also included below, where (a) shows GR_{5-20} from various condensation schemes, and (b) shows $\text{GR}_{1.9-5}$ from similar schemes. Note the inclusion of SVOC in GR_{5-20}), alongside the following statement

“Calculated growth rates according to the method of Nieminen et al. (2010) are presented in Figure S6 for both $\text{GR}_{1.9-5}$ and GR_{5-20} . Best agreement for GR_{5-20} is when condensation of SVOC,

LVOC, ELVOC, MSA, HIO₃ and H₂SO₄ is considered, and best agreement for GR_{1.9-5} is seen for condensation of all these except SVOC. The uncertainties in this method are large, and assumptions of irreversible condensation of SVOC onto particles of 5 nm likely lead to overestimations; however, these results confirm the essential role of the condensation of organic compounds to produce high growth rates observed in urban environments.”



Major Comment 4: Strong conclusions regarding pathway for NPF

Based on the results from the analysis, the results are not as conclusive as claimed both by title (“molecular insights”) and in the text (slightly different in abstract/discussion/conclusions). I would, therefore, suggest to rephrase the statements in text.

line 111: related to problem with title: “we examine molecular level evidence . . . at the critical diameter”- even the authors state “cluster identity cannot be identified” (line 276), so this wording is not appropriate.

line 293: It is clear that the results are similar to Kurten 2016 (except really high J). Mind that they conclude that SA-H₂O-DMA or other pathway could not be confirmed with their data.

lines 301-302: I suggest softening this conclusion. Also check abstract (lines 38-39) for consistency.

line 387: The first and main conclusion is that NPF “plausibly proceeds by the formation of clusters involving sulphuric acid and highly oxygenated organic molecules, with likely involvement of bases”. You clearly state earlier in line 276 that cluster identity cannot be directly identified and suggest that HOM is important for growth (not NPF, as far as I understood from text). So I am not sure what “plausibly” refers to. Please rephrase to reflect that results are not so conclusive, but the hypothesis is as follows (. . .).

Response 4: Our conclusions regarding the potential nucleation mechanisms in Barcelona have changed slightly given the updated parametrisations of H₂SO₄-DMA-H₂O nucleation from Kürten et al. (2018). We have made the following changes:

The wording at line 111 has been reworded to the following

“Here, we examine gas phase mass spectral evidence and particle formation rates at the critical diameter from sulphuric acid in Barcelona”

Further, at the end of section 3.2 we emphasise the following

“however, as no higher-order clusters were observed, we cannot establish this with certainty.”

In our conclusions we now make the following statements

“We show new particle formation rates in Barcelona are linearly dependent upon the sulphuric acid concentrations, and while formation rates far exceed that of H₂SO₄-BioOxOrg-H₂O nucleation, they fall short of those of H₂SO₄-DMA-H₂O nucleation at 278 K, as does the sulphuric acid dimer:monomer ratio, possibly explained by cluster evaporation due to high temperatures in summertime Barcelona (303 K during events), and limited pools of gas-phase amines. These results are similar to reports of nucleation rates in rural Germany (Kürten et al., 2016). As no higher-order clusters were directly measured, we cannot determine nucleation mechanisms with certainty, and an involvement of HOMs in nucleation is plausible.”

We feel that this better reflects the conclusions of this work, and have made similar changes to the abstract.

Minor Comments:

line 63: I appreciate the use of very recent references, but surely there are older references/review that appropriate to be referenced here (from other groups).

Response: This has been updated with reference to work across some diverse urban environments across the last 10 years, as well as a recent review from Lee et al., (2019), as this contains a section on NPF in polluted conditions with many recent references.

line 74: near-kinetic limit is achieved at certain DMA concentrations, can be specified here (Almeida et al. 2013).

Response: This section now reads

“Nucleation of sulphuric acid, DMA and water proceeds at, or near to the kinetic limit in a chamber at 278 K when DMA mixing ratios are sufficient (Almeida et al., 2013; Kürten et al., 2014).”

line 143: In Brean et al. 2019, calibration coefficient of 7e8 cm⁻³ was used. Were any further adjustments done in this work to get 3e9 cm⁻³?

Response: In this work, we correlated our uncalibrated data with a sulphuric acid proxy (Mikkonen et al., 2011), R²=0.49 and took the gradient as our calibration constant. We compare with (Brean et al., 2019) to ensure our figures are in the correct ballpark, accounting for differences due to different voltage tune. For clarity, this now reads

“Signals except for those of amines and ammonia are divided by the sum of reagent ion signals and multiplied by a calibration coefficient to produce a concentration. A calibration coefficient of 3×10⁹ cm⁻³ was established based upon comparison with a sulphuric acid proxy (Mikkonen et al.,

2011) and is in line with a prior calibration with our instrument (Brean et al., 2019). Uniform sensitivity between H_2SO_4 and all other species measured by CI-API-ToF apart from amines and ammonia was assumed in this work. This introduces some uncertainties, as it relies upon both collision rates and charging efficiencies to be the same within the ionisation source for all species. Amine and ammonia signals are normalised to the nitrate trimer signal (Simon et al., 2016). Prior reports of ammonia and amines as measured by CI-API-ToF employed corona discharge systems, which utilise higher concentrations of nitric acid, thus we report normalised signals.”

lines 216-217: Sulphuric acid time series/dial profile are discussed, which is not shown anywhere. Consider including time series of measured parameters in the Supplementary.

Response: We now include the diurnal profile of sulphuric acid as Figure S2

Figure 3: Was it possible to reliably fit amines and ammonia from the spectra? Kurten et al. 2016 mentions that at high RH, nitrate-water cluster interferes with the fitting of C2 amines. Please comment and preferably provide a figure of the peak fitting. Please also include Kurten et al. 2016 into references when talking about detecting ammonia with nitrate CI-API-TOF.

Response: We provide correlations of each amine as clustered with the dimer + trimer, and see a small RH% dependence on this ratio. The RH interference isn't too significant in this data, as the interference is seen, for example between the $(H_2O)_6(HNO_3)_{0-1}NO_3^-$ ions and the $(DMA)(HNO_3)_{1-2}NO_3^-$ ions. As our system is slightly different to that of Kürten et al. (2016), with lower concentrations of HNO_3 in the front end of the system, this H_2O clustering interference is smaller. We also provide intercorrelations of the amines with one another, and example peak fits of C₂ amines in Figure S1.

lines 234-236: You probably mean to say that TMA won't efficiently cluster with nitric acid and won't be detected. It is not very clear, please rephrase.

Response: We realise this is a mistake in the text and rephrase this whole section as follows

“Although high emission fluxes of TMA are expected in this environment, they are not present in our spectra, although this ion has been reported previously (Kürten et al., 2016).”

line 267 and Fig. 5: I think CLOUD data presented here is J1.7 for 278 degC, isn't it? Please mention it in the text and/or caption.

Response: This is correct. We specify this in the text as follows

“Data from these chamber experiments is for 278 K and 38 - 39 % relative humidity.”

And in our figure captions

“CLOUD chamber experiments were performed at 278 K and 38 – 39 % RH.”

We discuss the implications of these differing temperatures as in the response to major comment 1.

line 282-283: “Sulphuric acid dimer roughly represents the strength of sulphuric acid clustering in the nitrate CI-API-TOF”. - I am not sure what this means. Did you mean it tracks atmospheric clustering?

Response: This has been reworded for the sake of clarity also, and now reads

“The sulphuric acid dimer:monomer ratio is elevated by the presence of gas-phase bases such as DMA, and this elevation is dependent upon both the abundances and proton affinities of such bases

(Olenius et al., 2017). Upon charging, evaporation of water and bases from sulphuric acid clusters occurs, and thus these are detected at sulphuric acid dimer (Ortega et al., 2012, 2014)."

line 287: "lower limit" – do you mean upper? Please explain how IIC was calculated, or was it taken directly from Kurten et al. (2016)? I would like to note that the reaction time in drift tube in Kurten et al. (2016), is 50ms, what is it in your instrument?

Response: The reaction time in our Nitrate CI system is also 50 ms, so the calculated IIC as according to Zhao et al. (2010) is the same. "Lower limit" here was used to state that this is the lowest ratio that could be expected to be seen from this experimental setup, and has been reworded appropriately to

*"The dashed line represents the **ratio** that would be seen due to ion induced clustering (IIC) in the nitrate chemical ionisation system for a 50 ms reaction time (Zhao et al., 2010)"*

line 318: Is this a relevant reference for C10 HOMs that are coming with inland air mass?

Response: Our reference here is relevant to the discussion presented in Querol et al. (2017) regarding the inland air masses that flow towards the sea at night. This now reads as follows

"These nighttime HOMs will therefore mostly be derived from biogenic emissions which undergo more rapid nocturnal oxidation, and are likely transported from inland by the land breeze during night (Millán, 2014; Querol et al., 2017)."

lines 323-328: a series of unconnected sentences. please rephrase to make more logical.

"condensational growth being a reversible, step-wise kinetic process" – I am not sure what you mean, please rephrase. Do you mean "higher HIO₃ cannot account for stopped growth beyond 10nm?. rephrase it to fit in HOM discussion or move it to section 2.2.

Response: "higher HIO₃ cannot account for stopped growth beyond 10nm" was indeed the initial statement, but this sentence has now been absorbed into a more thorough discussion of particle growth.

In line 334, it is said that "precursors for these HOMs are presumed to be largely isoprene, alkylbenzenes and monoterpenes". However, in abstract (line 40) and in conclusions (line 394), isoprene is lost.

Response: These have been re-added to the abstract and conclusions

line 346: include also isoprene reference.

Response: We include a reference to Massoli et al., where isoprene HOMs are reported in the ambient environment.

line 369: you mention that HOM > 500 m/Q are increased during "full-events", however no data above 500m/Q is shown. Or do you mean HOM within 400-500 m/Q? This has to be clarified as this is also mentioned in conclusions. It can be shown as time series (or dial) of the sum of unit mass resolution peaks above 500 m/Q or box plot (I understand high masses are hard to fit/identify).

Response: These were repeated errors. This was meant to read >400 m/Q, and has been fixed.

line 389-390: "multiplicity of mechanisms has been shown to occur in chamber studies but has not been observed in the real atmosphere previously." - for instance Yan et al. (2018) showed two

regimes for NPF at SMEAR II station, Finland. Kurten et al. 2016 shows very similar results to this study and analogously (lines 390-391) points out high J, similarity to CLOUD for both SA-DMA-H₂O and SA-OxOrg studies. Please rephrase

Response: These sentences have been rephrased. At the end of section 3.2 we emphasise the following

“however, as no higher-order clusters were observed, we cannot establish this with certainty.”

Further, in our conclusions we have made the following changes

“while formation rates far exceed that of H₂SO₄-BioOxOrg-H₂O nucleation, they fall short of those of H₂SO₄-DMA-H₂O nucleation at 278 K, as does the sulphuric acid dimer:monomer ratio, possibly explained by cluster evaporation due to high temperatures in summertime Barcelona (303 K during events), and limited pools of gas-phase amines. These results are markedly similar to reports of nucleation rates in rural Germany (Kürten et al., 2016). As no higher-order clusters were directly measured, we cannot determine nucleation mechanisms with certainty, and an involvement of HOMs in nucleation is plausible”

We also include a reference to Yan et al. in the following sentence

“Conversely, research in remote boreal environments show that the mechanism of nucleation can modulate dependent upon the H₂SO₄:HOM ratio (Yan et al., 2018).”

Figure 7: Why HOM is shown in normalized signal, while in Fig. 3 it is in cm⁻³? How many days and in what time resolution are the data in this plot? Caption: “Influencing factors on VOC concentration” – on HOM concentration?

Response: We present an updated figure in the manuscript presenting HOM as molecules/cm⁻³ and have corrected the figure caption to now read

“Figure 7: Influencing factors on HOM concentration, showing (a) temperature plotted against C5-10 HOM signal, coloured by global radiation. Ellipsis shows 95% confidence on a multivariate t-distribution. (a) VOC concentration plotted against HOM signal. These are segregated by carbon number/VOC, i.e. C7 HOMs plotted against toluene, under the assumption that toluene oxidation is the main producer of C7 HOMs. Time for both plots is of hourly time resolution.”

Figure 8: Please provide a list of identified peaks in Supplementary and please explain what are the HOM observed at m/Q below 200 (quite high signals). Are they deprotonated species fitting with the “HOM” criteria?

Response: Peak list has been included as requested and referenced in the text as follows

“All ions identified are listed in Table S1.”

The inclusion of some small OVOCs as “HOM” was a printing error that has been fixed – these molecules are mostly deprotonated OVOCs such as dicarboxylic acids. The largest of these peaks belongs to malonic acid.

I suggest the authors would look into following and add more quantitative interpretation of results in the text.

-line 44: “significantly lower” – here easy to specify exactly how much, a factor of X.

-line 87-89: phrase “despite (extremely) high condensation sinks” repeats twice in a row. For comparison here, it would be nice what is meant by “extremely high”.

-line 100: please specify what is “frequent”, 40-lines 101-105, 195: “relatively high ozone”, “relatively low ozone (high compared with the rest of the year)”-can you rephrase or provide values as example? I also got lost in logic related to “relatively high” ozone during NPF and “maximum ozone episodes” with no NPF.

Response: Maximum ozone episodes refer to phenomena occurring in Barcelona where high particulate matter and ozone pollution occur concurrently, however, as this has little relevance to our discussion of NPF this has been removed. Our referenced sources here fail to provide a mean O₃ value through the year, but the cluster analysis results show elevations compared to the other most common clusters.

-line 244-245: “HOM . . . greatly enhanced”, “lower during burst”. Please provide values or a factor for HOM difference between days.

-line 248: “. . .HIO3 and . . .MSA are low”. I assume you refer to Fig S5d and e here? Relative to what are these concentrations low?

-line 270: instead of “extremely high” provide actual numbers from Fig. 5

-line 271: what does “broadly similar” mean?

-line 349: what does “broadly dependent” mean?

Response: We thank the reviewer for these very useful suggestions and have incorporated them through the whole text with the exception of one (see above).

Technical corrections

line 125: “interface” with small i

line 222: remove empty line

line 333: “temperature plotted against the signal of HOMs”. Isn’t it the other way around? Also see caption to Fig. 7.

line 368: insert: “enhancement” in comparison “to smaller alkylbenzene derived HOMs”.

line 401: repetition of line 394-395.

Figure 8. In-text (line 355), you mention that Fig. 8 has same days as Fig. 1. On Fig 1, there is 11/07/2018, while caption to Fig. 8 says 12/07/2018. Is it a typo? Please fix.

Response: We again thank the reviewer for pointing out these errors and have incorporated the appropriate corrections in the text.

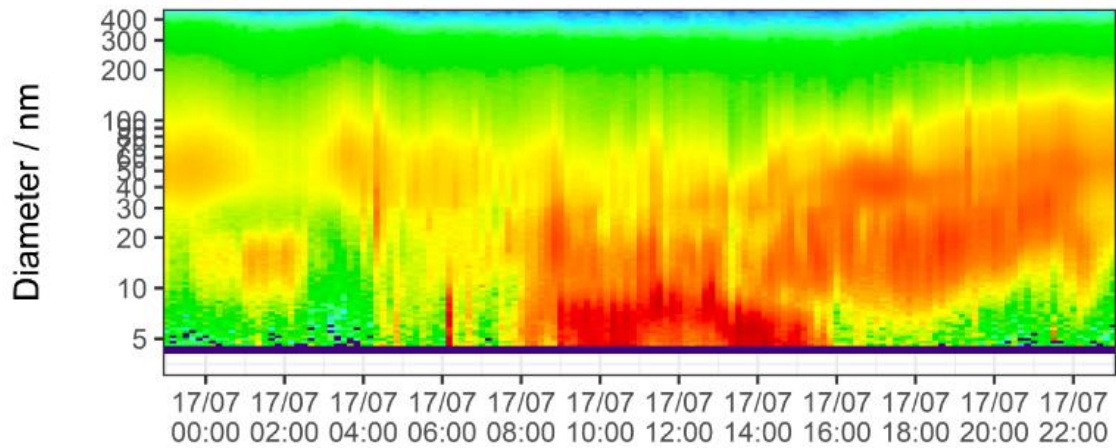
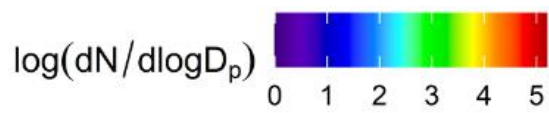
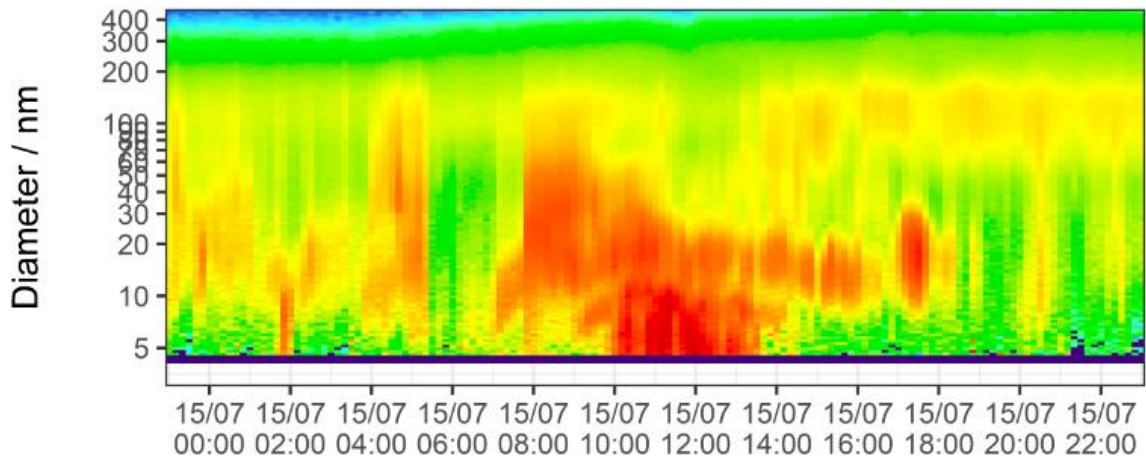
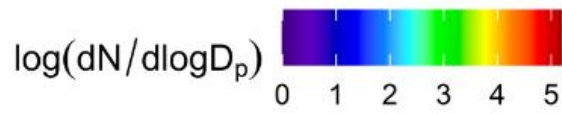
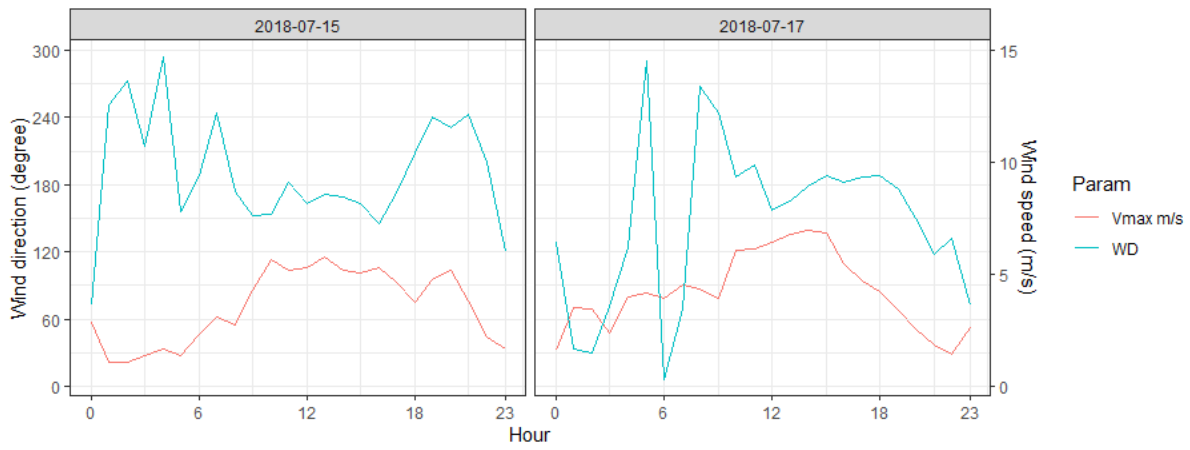
Referee 2

Major Comment 1: "Burst" events and air mass histories It is well established that NPF events for which formation followed by continuous growth over many hours can be observed (here called "full event"), are observable as such at a point site because NPF takes place more or less in sync across a fairly wide area (i.e. regional-scale NPF). Therefore, I wonder if the so-called "burst events" could be due to NPF that occurs more localized? I.e., that the more localized character of the NPF is why growth beyond 10 nm is not observed (with smallest observed particles having formed closest-by, whereas the largest observed new particles would have formed farthest away within the NPF area, upwind). The time resolution of the particle sizing measurements was quite low, but an estimate of the overall growth rate of the newly-formed particles could anyway be made, and from that, plus wind back trajectories, even the size of the hypothesized local NPF area could be estimated. Related to that: The manuscript mentions at least twice that NPF events, both "full" and "burst" types, are associated with "southerly and south-westerly air masses" (e.g. L 396-399). But Fig. S6 (showing back trajectories) clearly contradicts that statement! I suspect the mistake is in the text? That suspicion is (a) because I would expect a regional-scale BVOC-HOM-driven NPF event to have air masses NOT arriving from sea+town, and rather from inland, as suggested, presumably, in Fig. S6; and (b) upon noticing that the burst event shown in more detail (e.g. Figs. 1 and 8; July 15) is indeed the only event that does actually show a southerly back trajectory in Fig. S6, which puts the trajectory mostly over the ocean, with the exception of its final path across town (and, possibly critically, over shipping, as the authors also point out) towards the measurement site.

We discuss potential uncertainties arising from such local sources, following an extended discussion of methods used to derive J and GR . The data do not support the concept of burst events arising from an event of limited spatial extent. If they did, we would expect to see the disappearance of the mode associated with the burst and appearance of a new size distribution reflective of the new air mass. In our observations, the growth curve levels off – a mode of particles appears <10 nm and stays around that diameter, persisting for the course of hours. The argument of growth processes being limited is consistent with low concentrations of SVOC/LVOC/ELVOC on these days.

“The above calculations rely on the assumption of homogeneous air masses, and while air mass advection, as well as primary particle emissions can cause errors in estimations of temporal changes in particle count and diameter, the appearance and persistence of a new mode of particles across a period of several hours is typically indicative of a regional process.”

The discussion of south-westerly air masses was based upon an analysis of local wind direction + speed measurements. From polar plots it was inferred that air masses arising from the south were dominant during NPF, however, we find the air mass history from HYSPLIT to be more informative for our analyses, and thus these were included. The wind direction and velocity across the burst event days is shown below, where it can be seen that air masses are stable following the appearance of the new mode, and persist across the period of multiple hours, giving us confidence that this is not due to advection of a different air mass.



Major Comment 2: *Number of events and data points* Line 204 states that there was only 2 "burst" events observed. Then, it would be very useful to also learn how many "full" event were observed (and how many were non-event days)? However, with only 2 samples for the burst events, how can Figs. 2 and 3 include a boxplot for conditions during those two events? That may be down to the too sparse explanation of what data were used for these boxplots. E.g., data from which times were used in each event class (a vs b vs c)? Is the boxplot based on event-wise averages, or something else? But then again, Figs. 4+ suggest there is many more events of each type (judging from # of markers). A Fig. S6 again suggests again there's only been 5 events total!? I have some suspicions, but overall I am left rather confused. So, all that clearly should be presented more clearly, i.e. data quantity and usage in figures, resolved by event type.

Response 2: We answer these questions as below:

The periods of data that have been included have been clarified in the following sections

Figure captions for Figures 2 + 3 include the sentence

"[F]rom hourly data. Event days include data across the full event day."

And further, we address potential issues arising from producing box plots from a limited dataset by also including a complementary Figure S2, containing diurnal cycles of all the included parameters of Figures 2 and 3. Data coverage for some figures was a little misleading. Labels for Figures 4 – 6 have been amended for consistency, containing the phrase

"Data is for hourly averages across NPF periods, typically within the hours 08:00 – 16:00."

And in the start of section 3.1 we include the following

"Figure 1(b) shows a nucleation day with growth to larger sizes >10 nm, termed "full-event", showing the growth through the course of the day. These fulfil all the criteria of Dal Maso et al. (2005). 4 events of this type were observed with CI-APi-ToF data coverage. Figure 1(c) shows a day with nucleation occurring, but no growth past 10 nm. These days are referred to as "burst-event" days. Here, NPF is seen to occur, but particles fail to grow past the nucleation mode. 2 such events were seen in this data with CI-APi-ToF data coverage, and both are accompanied by a distinct mode appearing beforehand at ~20-40 nm."

We also present an updated Figure S5 containing the identity of each event day .

Major Comment 3: *Section 2.2: Was "sensitivity" applied to any detected compound? From the rest of the paper I conclude that it was applied to H₂SO₄ and HOMs, but not to ammonia and amines. Overall, some more detailed discussion of concentration quantifications based on CIMS signals would be interesting. See also next comment.*

Response 3: We expand our discussion of quantifications in the following sentences

"Signals except for those of amines and ammonia are divided by the sum of reagent ion signals and multiplied by a calibration coefficient to produce a concentration. A calibration coefficient of $3 \times 10^9 \text{ cm}^{-3}$ was established based upon comparison with a sulphuric acid proxy (Mikkonen et al., 2011) and is in line with a prior calibration with our instrument (Brean et al., 2019). Uniform sensitivity between H₂SO₄ and all other species measured by CI-APi-ToF bar amines and ammonia was assumed in this work. This introduces some uncertainties, as it relies upon both collision rates and charging efficiencies to be the same within the ionisation source for all species. Amine and ammonia signals are normalised to the nitrate trimer signal (Simon et al., 2016). Prior reports of ammonia and amines as measured by CI-APi-ToF employed corona discharge systems, which

utilise higher concentrations of nitric acid, thus we report normalised signals. We present correlations of each of these bases clustered with the nitrate dimer plotted against measurements with the nitrate trimer, as well as their intercorrelations and example peak fits across Figure S1. C₂ amines, C₄ amines and ammonia were the only molecules of this kind found in our mass spectra.”

Major Comment 4: *Ammonia and amines: Section 2.2: Are there any estimates on detection limits or sensitivity regarding ammonia and amines? (And which amines are expected to be detectable?) There is some discussion of amine detection late in Section 3.1. But it would be helpful if that more general instrumental aspects of their detection would already be at least mentioned in the Methods section. Section 3.1: Which amines were actually detected? I think that information is only provided in the caption of Fig. 3 and in Fig. S1 (implying C₂H₇N and C₄H₁₁N), and that should be mentioned also in the main text, and much earlier. (And is that in agreement with expectations from amine abundances and amine-specific sensitivities indicated by previous studies (if any)?) The authors do discuss relative sensitivities later in Section 3.1, implying that sensitivities to all those small bases correlate to NPF enhancement potential. That argument could be brought sooner. Figure 3, and related discussion (Section 3.1, page 9): From the beginning of the discussion of Fig. 3b, I have been wondering why are ammonia & amine signals only presented as a sum of all signals? Even if the concentrations of these compounds could not be quantified in these measurements, it still appears it would be more insightful not to lump them together. Especially given the respective differences in NPF enhancement capabilities. Also, if lumped together, there is the chance that differences in atmospheric abundance of ammonia vs. amines (almost certain, as also described in Section 3.1) or in instrument sensitivity to these compounds (plausible) lead to one compound (or group of compounds) controlling the lumped signal... Only very late in Section 3.1, L 241 mentions that all the bases’ time series correlated quite well, thus justifying the lumping. That argument for lumping should be made more clearly and earlier. And supported by a figure or two in the Supplement. L 217: From looking at Figure 3, I am not sure I would also arrive at the conclusion that ammonia+amine concentrations were enhanced for full events. The medians are practically identical. Have the authors applied some quantitative measure of statistical significance regarding differences of the parameters presented in Figs. 2-3?*

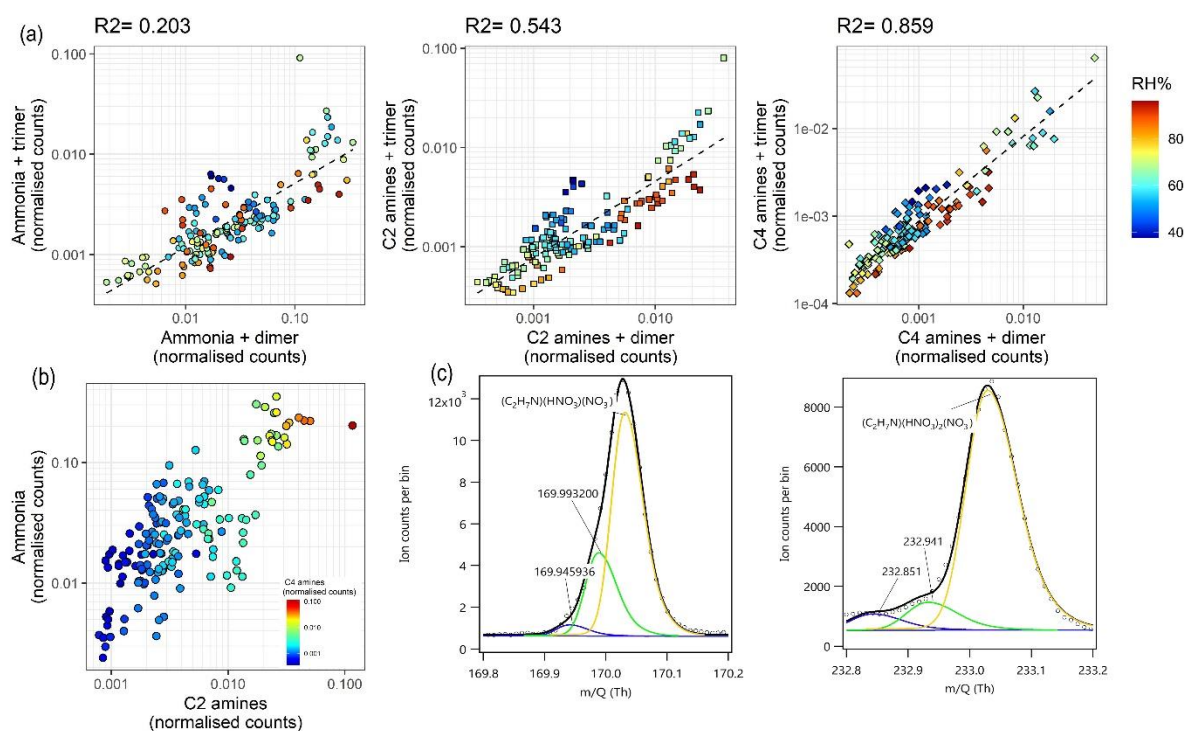
Response 4: Unfortunately, the two prior reports of amines (Simon et al., 2016) and both ammonia and amines (Kürten et al., 2016) from NO₃⁻ ionisation used a corona discharge CI-APi-ToF system, which is distinct from our own for these purposes as the amount of reagent ion in the front end of the instrument is higher. The result of this is that detection limits and sensitivities are likely significantly different to our own. We can apply the sensitivities of Simon et al. (2016), where a sensitivity a factor of 10 different from that for H₂SO₄ is applied, and this gives us a mean C₂ amine concentration of just under 10 pptv across the campaign.

Amines and ammonia have been considered as one “sum” here as they follow similar trends. Below we tabulate the Pearson correlation coefficients between ammonia and amines, as well as a plot of ammonia plotted against C₂ amine signals, coloured by C₄ amine signals. We also include all of this information in Figure S1, which is included below (see panel (b)) as well as an expanded Methods section (see quoted text above) to clear issues of placement of this information within the manuscript, while the following is included in the text.

“We present correlations of each of these bases clustered with the nitrate dimer plotted against measurements with the nitrate trimer, as well as their intercorrelations and example peak fits across Figure S1”

	C ₂ am.	C ₄ am.
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NH_3	0.71	0.58
$C_2 am$		0.94



We also present individual diurnals for these compounds in a new Figure S2. Fortunately, the magnitude of each signal of each of these compounds is similar, and thus our Figure 2 is not weighted towards ammonia. This is not as significant as the difference expected in typical atmospheric mixing ratios (Ge et al., 2011). We realise this is evidently a function of an insensitivity towards ammonia. We remove ammonia from this figure for both this reason, and because the presence of alkylamines even in low mixing ratios will likely be more significant for nucleation processes than ammonia (Almeida et al., 2013; Yao et al., 2018) as even modest mixing ratios of alkylamines can substitute ammonia in sulphuric acid-ammonia clusters (Kupiainen et al., 2012), and thus argue that these are the important aspect of these measurements. The argument of significance was made based upon a simple t-test.

Major Comment 5: Figure 5 and/or related discussion Environmental conditions should be mentioned and discussed, in particular in light of the CLOUD experiments being compared to: T , RH , compound concentrations (as known), as all those will have affect formation rates. Also: given available parametrizations of NPF rates as a function of sulfuric acid, HOM (or "BioOxOrg") and amine/ammonia concentrations, are the authors able to explain observed NPF rates and infer concentrations of either involved BioOxOrg or amines? Combined with expected amine concentrations and measured VOC and HOM concentrations, some closure could be attempted. Some hand waving may be necessary, but the attempt could be quite interesting. And it could strengthen the paper, especially if it can be argued that closure \sim works out. Similarly, growth rates of new particles could be estimated (see also 1st comment) and evaluated against HOM concentrations.

Response 5: Regarding the temperature and relative humidity dependencies, we include the following

“Model studies of sulphuric acid-amine nucleation show a decline in nucleation rate with temperature (Almeida et al., 2013; Olenius et al., 2017), as the evaporation rate of sulphuric acid-amine clusters will increase with temperature (Paasonen et al., 2012). Conversely, evaporation rates of such small clusters, and resultant nucleation rates tend to increase modestly with increases in relative humidity, most pronounced at lower amine concentrations (Almeida et al., 2013; Paasonen et al., 2012). Despite this, high nucleation rates at temperatures nearing 300 K have been reported previously (Kuang et al., 2008; Kürten et al., 2016), although these tend to show a temperature dependence (Yu et al., 2016).”

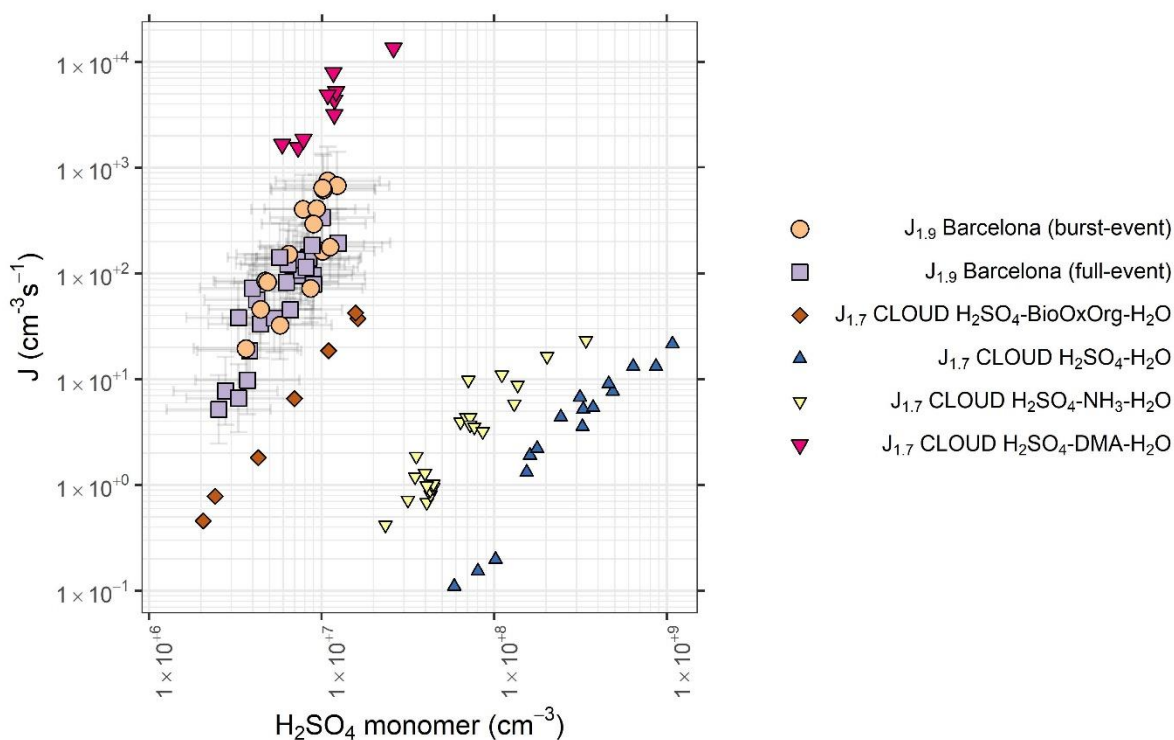
We also emphasise the differing conditions of these chamber results in our figure captions. Included in the new version of the manuscript is an updated parametrisation of H₂SO₄-DMA-H₂O nucleation from Kürten et al. (2018), which somewhat changes our conclusions. The results of Kürten et al. (2018) involve re-analysis of the CLOUD data and produce nucleation rates around an order of magnitude higher than in Almeida et al. (2013). Thus, we produce a new Figure 5 included below. This produces a far more consistent picture, and we now suggest a nucleation pathway possibly involving sulphuric acid and amines, with rates reduced by around an order of magnitude by elevated temperatures, consistent with prior evidence (Olenius et al., 2017). We therefore include the following sections

“Particle formation plausibly operates by sulphuric acid-amine nucleation involving the measured C2 and C4 amines in our data, with nucleation rates hindered relative to those measured in the CLOUD experiments by elevated temperatures, and a decline to the sulphuric acid dimer:monomer ratio indicates that base concentrations may be limited. We cannot rule out an involvement of HOMs in particle formation processes, and, as no higher-order clusters were observed, we cannot establish sulphuric acid-amine nucleation with certainty.”

And conclude

“while formation rates far exceed that of H₂SO₄-BioOxOrg-H₂O nucleation, they fall short of those of H₂SO₄-DMA-H₂O nucleation at 278 K, as does the sulphuric acid dimer:monomer ratio, possibly explained by cluster evaporation due to high temperatures in summertime Barcelona (303 K during events), and limited pools of gas-phase amines. These results are similar to reports of nucleation rates in rural Germany (Kürten et al., 2016). As no higher-order clusters were directly measured, we cannot determine nucleation mechanisms with certainty, and an involvement of HOMs in nucleation is plausible.”

We would argue that the necessary hand waving might be too great to justify such a parametrisation as suggested above. An example of such a parametrisation would be that of Hanson et al., (2017), wherein a simple equation to derive J during H₂SO₄-DMA-H₂O nucleation is proposed, however, this requires a quantified DMA mixing ratio, and to assume one would introduce quite some uncertainties into this. However, we hope the reviewer find our updated discussion of potential mechanisms of nucleation compelling.



Major Comment 6: Figure 6 and related discussion (in particular L 289-293): Could the SA dimer signal also be affected by instrument settings? It is conceivable that some instruments or settings would fragment a certain fraction of SA dimer ions (that is, at some point after their formation by NO_3 -ionization), and that that fraction is instrument- or tuning specific. If so, conclusions can likely still be made anyway from comparisons between measurements within the same campaign (provided settings remained the same), and probably even from comparisons across campaigns/longer time periods as long as the same instrument was used. But could such instrumental differences cause (part of) the discussed discrepancies between dimer signals here and results shown from other field and lab experiments? My feeling is that the SA dimer anion is stable enough that such instrumental fragmentation should not be expected, but I would ask the authors to at least point that out (i.e., if my "feeling" can be defended based on previous studies – I apologize for not remembering expected cluster stabilities vs the fragmentation potency of the API-TOF instrument), OR recognize/discuss potential issues.

Response 6: The region of the instrument most likely to produce fragmentation of the sulphuric acid dimer is the collision dissociation chamber (CDC). A factor of 1.2 difference in sulphuric acid dimer concentration compared to that expected from ion induced clustering in the chemical ionization source has been shown in a chamber study (Kürten et al., 2015). This is consistent with the high binding energies of the $\text{H}_2\text{SO}_4\text{HSO}_4^-$ cluster, in excess of 40 kcal mol^{-1} (Curtius et al., 2001), around a factor of two higher than other commonly measured ions such as HOM-HSO_4^- or HOM-NO_3^- clusters (Froyd and Lovejoy, 2012; Kurtén et al., 2007; Zanca et al., 2020). The result of this is an evaporation rate of around $2.7 \cdot 10^{-15} \text{ s}^{-1}$ compared to an evaporation rate of DMA out of the $(\text{DMA})_2\text{H}_2\text{SO}_4\text{HSO}_4^-$ ion of $2.7 \cdot 10^{-1} \text{ s}^{-1}$. Further evidence is that when a negatively charged sulphuric acid trimer makes its way into the instrument it is fragmented largely into a charged sulphuric acid dimer, and a neutral monomer (Passananti et al., 2019). To our knowledge there have been no results published on the effect of instrument tuning on the sulphuric acid dimer (just those on the sulphuric acid trimer and HOMs as referenced above). We do, however, include the

following sentence in our manuscript (with further mention of the potential effect of condensation sinks

“The binding energy of the bisulphate- H_2SO_4 ion is in excess of 40 kcal mol^{-1} (Curtius et al., 2001), and thus minimal declustering of the dimer is expected within the CI-APi-ToF instrument – however, declustering of higher order sulphuric acid clusters has been shown to be sensitive to voltage tune (Passanati et al., 2019), and this likely extends to the dimer also, and as such discrepancies between sets of results due to instrument setup cannot be ruled out. The ratio of sulphuric acid dimer:monomer is also highly sensitive to condensation sinks, with a difference in dimer concentration of approximately a factor of 4 expected at 10^7 cm^{-3} between 0.001 s^{-1} (a clean environment) and 0.03 s^{-1} (our calculated condensation sinks during NPF events) (Yao et al., 2018) and thus our low dimer:monomer ratio can, in part, be explained by elevated condensation sinks.”

Minor comments:

Abstract: a) I would already here mention roughly the most important measurement methods used for the study. No details, and a side sentence may be enough. Especially as there is talk about sulfuric acid monomer and dimer concentrations, I would have liked to be informed already here that those are based on NO_3 -CIMS measurements.

Response: We have included the following in the abstract

*“Here, we study summertime NPF in urban Barcelona in NE Spain **utilising particle counting instruments down to 1.9 nm and Nitrate CI-APi-ToF.**”*

b) Growth beyond 10 nm: my immediate thought already here was that limited growth could also be a sign of more local NPF (vs. regional, i.e. on a larger geographical scale), in which case the observed lower concentrations of low-volatility organics could be irrelevant (see 1st major comment). So I would already here, in a compact way, give the reasoning for claimed conclusion.

Response: This has been included.

Introduction: The last paragraph (review of NPF observations in the Barcelona area) is a somewhat confusing to read. Should be restructured for clarity.

Response: This has been restructured – we have removed the mention of high ozone events as they have little relevance to NPF, as well as mention of ozone across other urban environments. This section now reads

“Urban Barcelona sees frequent, strong summer-time NPF events occurring on 28% of days. These events are associated with high insolation and elevated ozone ($\sim 60 \mu\text{g m}^{-3}$) when considering the whole year (Brines et al., 2014, 2015). Ground-level observations report NPF events starting typically at midday, and either occurring in urban Barcelona and the surrounding regional background simultaneously, or isolated to either urban Barcelona or just the regional background (Dall’Osto et al., 2013). Vertical profiles over urban Barcelona reveal that NPF occurs at higher altitudes, and starts earlier in the day, as at a given altitude these events are not suppressed by early traffic peaks contributing to particle load (Minguillón et al., 2015).”

I found the usage of the term “background” not clear (end of introduction, beginning of methods, beginning of results).

Response: The inclusion of the term “urban background” refers to the urban environment over the scale of several km^2 . An “urban background site” is an urban site away from direct emission sources that is descriptive of a well mixed urban region. “Regional background” refers to Barcelona

and its surrounding regions, here on the scale of tens of kilometres. This terminology is well accepted in the air pollution literature.

L 143: The meaning of "sensitivity" here (in previous studies called also "sensitivity coefficient" or "calibration factor") remains a mystery for any reader who is not fairly intimate with operation of that CIMS instrument. So, I would at least cite some paper where that meaning is discussed, e.g. Kürten et al., 2012 (10.1021/jp212123n). Regarding the "sensitivity" value used, the authors cite here a previous calibration, paper also led by UBirmingham. However, I could not find, in that paper, how that calibration was performed. Indeed that paper contains the statement (in their section 2.2) "No sensitivity calibration was performed for these measurement..." Bottom line is that I have remained wondering where the used value ($3 \times 10^9 \text{ cm}^{-3}$) derives from. (Or simply from comparison to the sulfuric acid proxy?)

Response: To answer, and provide more detail re: calibrations, sensitivities etc., we include the following

"Signals except for those of amines and ammonia are divided by the sum of reagent ion signals and multiplied by a calibration coefficient to produce a concentration. A calibration coefficient of $3 \times 10^9 \text{ cm}^{-3}$ was established based upon comparison with a sulphuric acid proxy (Mikkonen et al., 2011) and is in line with a prior calibration with our instrument (Brean et al., 2019). Uniform sensitivity between H_2SO_4 and all other species measured by CI-API-ToF apart from amines and ammonia was assumed in this work. This introduces some uncertainties, as it relies upon both collision rates and charging efficiencies to be the same within the ionisation source for all species. Amine and ammonia signals are normalised to the nitrate trimer signal (Simon et al., 2016). Prior reports of ammonia and amines as measured by CI-API-ToF employed corona discharge systems, which utilise higher concentrations of nitric acid, thus we report normalised signals. We present correlations of each of these bases clustered with the nitrate dimer plotted against measurements with the nitrate trimer, as well as their intercorrelations and example peak fits across Figure S1."

L 221: It is not clear if this is a general statement regarding amine concentrations, or specific to observations (which, however, I understand could not be quantified, so I assume it's the former)? Should be clarified.

Response: This is indeed a general statement, we have included a reference to an appropriate source here.

Figure 1 is lacking all labels for the axes. At least the somewhat-less-obvious vertical axis should be labelled.

Response: Figure 1 has been amended (and individual days have been replaced with the average contour for each of these day-classes).

L 244: Should start new paragraph when starting discussion HOMs.

Response: Done

L244-245: It appears odd to discuss HOMs (Fig. 3b) with only 1 sentence, following a page of discussion of ammonia/amines! Should at least add a reference to a later section, where organics, including HOM, are being discussed.

Response: This section now reads

"HOM concentrations were greatly enhanced during full-event periods (factor of 1.5 higher compared to non-NPF mean), but lower during burst-event periods (factor of 1.2 lower compared

to non-NPF mean), implying their necessity for growth. The sources and implications of these HOMs are discussed in section 3.3”

L247-248: This last sentence of Section 3.1 is very vague. How high (or low) are those concentrations of marine compounds? (Or estimated to be?) What observations would have been considered evidence FOR an influence of oceanic emissions on NPF/growth?

Response: These have been stated and this section now reads

“concentrations of iodine and DMS-derived acids such as iodic acid (HIO_3) and methanesulphonic acid (MSA) are low ($7.8 \cdot 10^5$ and $3.3 \cdot 10^5 \text{ cm}^{-3}$ respectively), indicating a small influence of oceanic emissions on particle nucleation/growth”

Evidence for the influence of oceanic emissions would, in this context, be taken as the presence of high mixing ratios of oceanic-emission derived acids. An example would be a high MSA: H_2SO_4 ratio as both arise plentifully from DMS oxidation (Hoffmann et al., 2016), or high concentrations of iodine acids derived from biogenic iodine vapours (Sipilä et al., 2016).

Fig. 4: How steep ist that slope? That is usually interesting information, at least for comparing with other studies.

Response: This has been included in the figure caption. The value of the slope is $4.9 \cdot 10^{-5} \text{ s}^{-1}$

L 261: Does "losses" refer to amines? If so, I don't see how photochemical reductions of amine mixing ratios could mask an actual dependence of NPF rates on those mixing ratios. (If I caught the inference correctly.)

Response: That was indeed the inference, but this point is secondary to the more important one made before. This section now reads

“In the example of alkylamines, their gas phase concentration may decrease due to clustering with elevated sulphuric acid, as they cluster at around a 1:1 ratio at high amine mixing ratios (Kürten et al., 2014) (and therefore they will not be detectable as free amines). Further, if amines are present at a few pptv, their mixing ratios are significantly higher than our ambient measured sulphuric acid concentrations, and will be sufficient to accelerate nucleation rates (Almeida et al., 2013). Photochemical losses will also be greater during the periods of highest NPF rate (Ge et al., 2011b).”

L 273: Suggest rephrasing to make it clear immediately that the discussion shifts from literature results to new results. And it is not clear which observations the last part of the sentence refers to ("on these days").

Response: We have clarified this. It reads

“indicating similar mechanisms of formation, despite lower HOM concentrations on burst-event days”

L 282-283: ambiguous what is meant by "strength of sulfuric acid clustering". It becomes clear thereafter, but if I understand correctly, "strength" is not the right word.

Response: This has been changed and now reads as follows

“To further explore the relationship between sulphuric acid clusters and the rate of nucleation, the sulphuric acid dimer:monomer ratio is plotted in Figure 6. The sulphuric acid dimer:monomer ratio is elevated by the presence of gas-phase bases such as DMA, and this elevation is dependent

upon both the abundances and proton affinities of such bases (Olenius et al., 2017). Upon charging, evaporation of water and bases from sulphuric acid clusters occurs, and thus these are detected as sulphuric acid dimer (Ortega et al., 2012, 2014)."

L 293+: Could the point raised in major comment (6) explain the flatter slope observed here vs. the Germany observations?

Response: We include a discussion of this, alongside a discussion of the result of elevated condensation sinks as mentioned in the response to major comment 6

L 337: Getting confused here. Should it read "not largely radiation dependent" instead of "not largely temperature dependent"?

Response: This section has been modified to read as follows

"A lack of correlation between other VOCs and their HOMs confirms that this relationship between HOMs and temperature is not a function of enhanced VOC emission fluxes from, for example, evaporation, except in the instance of isoprene."

L 342: Should they be transport FROM inland by the land breeze?

Response: Correct. This has been amended.

L 357, 361: Should be explained what is meant by "detailed criteria", and by "updated criteria" (i.e. what are the respective criteria).

Response: We update this section to the following, and move it slightly earlier

"Oxygenated volatile organic compounds (OVOC) are defined as species visible in the nitrate CI-API-ToF that do not classify as HOM. Here, the first of the three criteria provided by Bianchi et al. (2019), that HOM must be formed by peroxy radical autoxidation, cannot be applied to define HOM, as knowledge as to whether a molecule is a result of autoxidation requires sound knowledge of the structure of the precursors, oxidants and peroxy radical terminators present, however, the number of molecules observed with $nN = 2$ is around an order of magnitude lower than that for $nN = 1$, where the primary source of multiple nitrogen functionalities would be multiple peroxy radical termination reactions from NO_x , and we argue that while multiple generations of oxidation have been shown to occur in aromatics (Garmash et al., 2020), it is a small contributor to the concentration of what we class as HOM. The second criterion to define HOM are that they must be formed in the gas phase under atmospherically relevant conditions, which we deem appropriately fulfilled as all CI-API-ToF measurements are of gas phase compounds, and the final criterion is that HOM must contain more than 6 oxygen atoms. To attempt to satisfy these criteria as best possible, the criteria of both containing 6 oxygen atoms and 5 carbon atoms or greater and having an O:C ratio >0.6 is applied, as these molecules will all plausibly fulfil the updated criteria of "HOM"."

L 367: could be informative to point out some of those formulae explicitly

Response: We include the following

"The full-event day sees enhancements to smaller OVOCs and HOMs compared to the non-event day, especially around 150-200 m/Q, which contains peaks corresponding to dicarboxylic acids and isoprene oxidation products. Some of the largest peaks in the mass spectra correspond to formulae seen arising from the enhanced OH oxidation of alkylbenzenes (such as $C_7H_7NO_6$) (Molteni et al., 2018; Wang et al., 2017)"

L 371-372: would be instructive to be more specific regarding "large" and "smaller"

Response: We include the following

“During full-event periods, these peaks are both more numerous and larger, with a factor of two difference in total peak area in this m/Q range”

L 378-382: Something went wrong with this sentence, especially the first part. Think I get the idea, but not sure.

Response: We edit these sentences to now read

“These elevations to condensable OVOC and HOMs on particle formation days with growth are consistent with particle composition data as measured by ACSM (Figure S9). Particle composition on full-event days shows an elevation to organic mass concentration in the late evening and night around when new particles from NPF will reach sizes detectable by the ACSM (~75 nm, Ng et al., 2011). Organic mass between 16:00 – 23:00 is 3.5 $\mu\text{g m}^{-3}$ on burst-event days, versus 7.8 $\mu\text{g m}^{-3}$ on full-event days.”

L 381: Please state the size range the ACSM is sensitive to

Response: This is stated in the methods now

“Continual monitoring of composition and mass of submicron aerosol >75 nm was carried out with an Aerosol Chemical Speciation Monitor (ACSM, Aerodyne, USA) (Ng et al., 2011).”

L 393: Could the mechanisms also support each other (i.e. be combined) rather than be in competition?

Response: This statement in the conclusions has been softened, and now reads

“[F]rom our evidence, we hypothesise a mechanism proceeding by the formation of clusters involving sulphuric acid, with potential involvement of both HOMs and amines, as we measure both small alkylamines, and molecules of class ELVOC, both of which have shown themselves capable of forming particles in conjunction with sulphuric acid. This potential multiplicity of mechanisms has been shown to occur in chamber studies and is markedly similar to reports from rural Germany (Kürten et al., 2016).”

Fig. S6: Please include information on which kind of events are shown, and when (... see also major comments (1) and (2))

Response: This figure has been updated with the appropriate information

Technical comments

L 61: i.e. should be e.g.?

L 79/80: those "loss processes" haven't been mentioned yet, so would be instructive to name the most important ones (here, for < 50 nm particles)

L 81: "at these diameters" ... rather "from these diameters onward"

L 86: I believe that range should read "0.5-11" (i.e., 11 instead of 1.1)

L 86-90: two redundant consecutive sentences. combine.

L 138: if I remember correctly, the flow containing the reagent ions is not "guided into the sample flow", but rather only the ions are guided there (electrically).

L 162: redundant mentioning of "4 flows"

Fig. 3, a and b: the "+" in the exponents (tick labels) are conventionally omitted

L 371: compromise -> comprise

Figs. S6 and S7 are not referred to in the main text (not sure if that's a problem)

Response: We thank the reviewer for these comments – these have been amended in the text.

Referee 3

Hi Josef, thank you for the insightful comments. We have separated these into individual comments and addressed these in-turn below

Comment 1: Since molecular level knowledge of new particle formation in urban areas is still very poor I read this paper with high interest. I strongly support the reviewers asking for a better description of the analysis methods and more quantitative information. For example in Figure 1 b no distinct particle evolution is seen. This looks more like an advection of an air mass. It would be worthwhile to show in an example how the nucleation rate was determined? The same applies to the growth rate calculation. GRs seem to be quite high but time resolution of measurements rather low for such events. For example, the time resolution of the PSM measurements is 10 min. Thus for $GR > 6$ nm/h no growth rate in the sub-3nm range can be determined.

Response 1: In line with these suggestions, we have produced an expanded methods section. We use a time-delay method utilising the 1.9 nm size cut of the PSM and compare it to the integrated particle counts from the NanoSMPS + SMPS measurements, giving the time-delay between measuring particles at 1.9 and 5 nm, thus, particle growth rates would need to reach ~ 18 nm h⁻¹ in order to escape the resolution of our measurements. We realise this has its potential uncertainties, and include a discussion of such uncertainties in our results section. Alongside there is an expanded Methods section where we state the following:

“The formation rate of new particles at size d_p is calculated as follows:

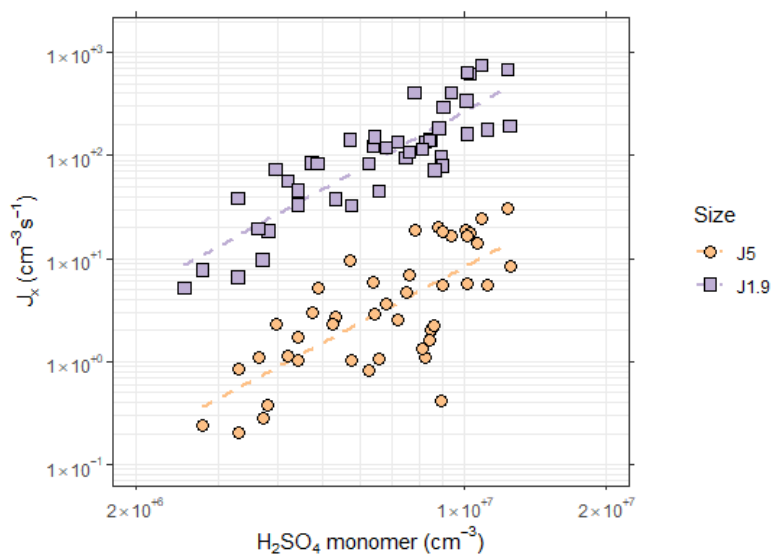
$$J_{d_p} = \frac{dN_{d_p}}{dt} + CoagS_{d_p} \cdot N_{d_p} + \frac{GR}{\Delta d_p} \cdot N_{d_p} \quad (2)$$

Where the first term on the right-hand side comprises the rate at which particles enter the size d_p , and the latter two terms represent losses from this size by coagulation and growth respectively. J_5 was calculated using the data between 5-10 nm, and $J_{1.9}$ was calculated using the measurements between 1.9 – 4.5 nm. We also calculated $J_{1.9}$ from our NanoSMPS data, employing the equations of Lehtinen et al. (2007). $J_{1.9}$ from both methods showed reasonable agreement ($R^2 = 0.34$) Agreement between J_5 and $J_{1.9}$ for each method was similar ($R^2 = 0.37$ and $R^2 = 0.38$ for $J_{1.9}$ calculated from PSM data and from Lehtinen et al. (2019) respectively). See Kulmala et al. (2001) for more information on calculation of coagulation sinks and formation rates. Growth rates between 4.5 – 20 nm were calculated according to the lognormal distribution function method (Kulmala et al., 2012), whereas those between 1.9 and 4.5 nm were calculated from PSM data using a time-delay method between PSM and NanoSMPS data. Systematic uncertainties on our calculated $J_{1.9}$ values include 25% method uncertainty (Yli-Juuti et al., 2017), with a further 25% arising from uncertainties in PSM cutoff (± 0.5 nm), as well as a 10% uncertainty in counting errors. A 50% error arising from calculated coagulation sink is also applied (Kurten et al., 2016). The above calculations rely on the

assumption of homogeneous air masses, and while air mass advection, as well as primary particle emissions can cause errors in estimations of temporal changes in particle count and diameter, the appearance and persistence of a new mode of particles across a period of several hours is typically indicative of a regional process.”

Comment 2: *It is hard to believe that J5 and J1.5 are almost similar. At H₂SO₄ = 2E06 the reported J5 is even ten times higher than J1.5.*

Response 2: We confirm that the initial Figure 4 was a printing error – in the updated form of this manuscript we present an updated Figure 4 using J_{1.9} values. For your consideration below we also include J₅ and J_{1.9} on the same plot which differ by around the expected order of magnitude or so, given the high coagulation sinks.



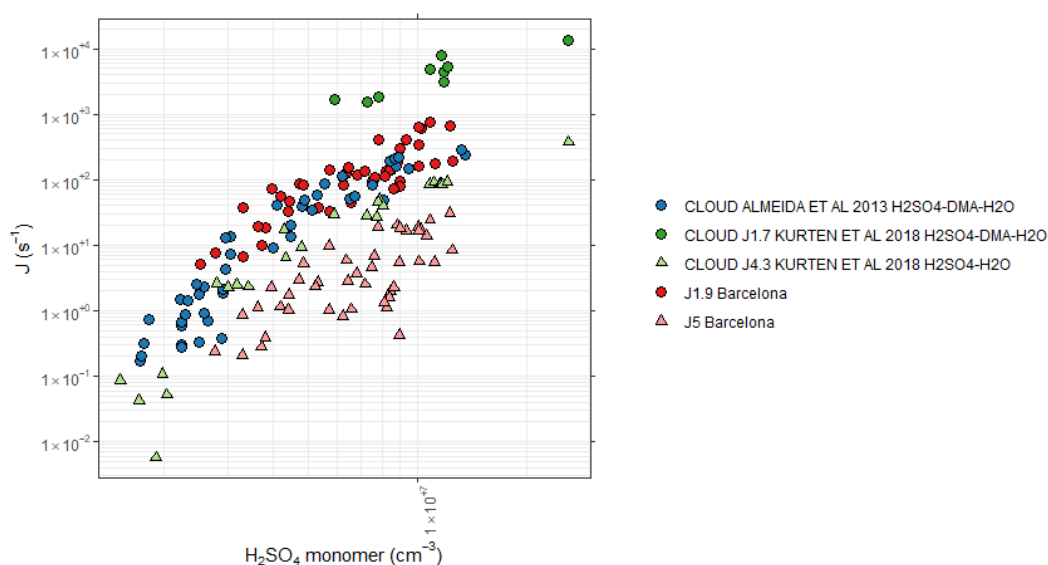
Comment 3: *What are the uncertainties of J1.5, J5 and GR at different sizes? As pointed out by Kulmala et al. (2017) nucleation and growth rate are connected. It would be very helpful for the reader if the authors would also report the growth rates, as they have been determined anyhow. Kulmala et al. (2017) also show that the survival probability of small clusters becomes very low at such high coagulation/condensation sinks as reported in this study. A discussion of this phenomenon should be included.*

Response 3: The revised manuscript now contains stated growth rates as calculated from our particle count data, as well as a mean survival parameter. We include these in the following sentences

“The survival parameter (P) as suggested by Kulmala et al. (2017) is defined as $CS \cdot 10^{-4} / GR$, and for this data is equal to 82, higher than other European cities. The occurrence of such a high P value should, in theory, inhibit the occurrence of NPF, but we show events happen readily under such conditions, akin to other heavily polluted megacities.”

We also include the updated J_{1.7} values of Kürten et al., (2018) in a new Figure 5, while parameterising our errors as follows: We estimate a 25 % error in GR arising from methods chosen (Yli-Juuti et al., 2017) and a further 25% arising from potential uncertainties in the PSM cutoff of ± 0.5 nm. A 10 % uncertainty has been presumed to arise from errors in counting, and a 50 % error arising from systematic errors in calculating coagulation sinks (Kürten et al., 2016). Similarly, we

present a +100% / -50 % error on sulphuric acid condensation arising from systematic errors in our estimation of C . We include a discussion of such errors in our methods, and further reference these uncertainties in the respective figure captions. Below, we show a comparison of our own $J_{1.9}$ and J_5 values with the CLOUD data from Almeida et al., (2013) and Kürten et al., (2018).



Comment 4: The authors strongly stress the role of organics in NPF. Measurements of the HOMs and their chemical composition are available. Thus, the authors could estimate if the concentration of very low volatility HOMs is high enough to account for the growth of few nm-sized particles. Such an analysis would support their conclusions on what drives NPF. In Figure 8 it is hard to see differences in the mass spectra at $m/z > 250$. Thus, most differences are at lower m/z and for compounds with high mass defect, that is low oxygen content, and thus high volatility.

Response 4: Regarding the interesting comment on growth rates, we have calculated $\text{Log}_{10}(C^*)(300\text{ K})$ as according to the simple linear equation provided in Bianchi et al., (2019), an updated form of that first presented in Donahue et al., (2011). As per their recommendations, we include an additional term to account for nitrogen functionality where every -O-NO₂ group reduces $\text{Log}_{10}(C^*)$ by 2.5. The volatility distribution of products is presented below, and is now included in the text of the paper. We discuss this volatility parametrisation and the implications in the following paragraphs

“DBE as calculated by equation 3 is equal to the number of pi bonds and rings within a molecule. Benzene, toluene, and similar aromatics have DBE = 4, naphthalene = 7 and monoterpenes = 3. DBE can be used as an indicator of sources when considering HOM in bulk. Saturation mass concentration as calculated by equation 4 can help describe capacity of a molecule to both condense onto newly formed particles and participate in nucleation. Figure 8 shows concentrations of HOMs and other oxygenated organic molecules binned to the nearest integer $\text{Log}_{10}(C^)(300\text{ K})$, coloured by DBE. Most measured molecules fall into the SVOC class ($0.3 < C^*(300\text{ K}) < 300\ \mu\text{g m}^{-3}$) which will mostly exist in equilibrium between gas and particle phase. High SVOC concentrations arise from fingerprint molecules for isoprene oxidation under high NO_x ($\text{C}_5\text{H}_{10}\text{N}_2\text{O}_8$) (Brean et al., 2019), and oxidation of small alkylbenzenes ($\text{C}_7\text{H}_8\text{O}_5$, $\text{C}_8\text{H}_{10}\text{O}_5$). LVOC and ELVOC ($3 \cdot 10^{-5} < C^* < 0.3\ \mu\text{g m}^{-3}$ and $3 \cdot 10^{-9} < C^*(300\text{ K}) < 3 \cdot 10^{-5}\ \mu\text{g m}^{-3}$ respectively) have greater contribution from molecules with higher DBE, i.e., $\text{C}_{10}\text{H}_{10}\text{O}_8$ arising most likely from PAH*

oxidation (Molteni et al., 2018), and C₁₀H₁₅O₇N, a common molecule arising from monoterpene oxidation in the presence of NO_x. The contribution of molecules with carbon number ≤ 9 to these LVOC is modest, and ELVOCs are entirely comprised of molecules with carbon numbers ≥ 10 and DBEs of 8 and 4. No molecules classed as ultra-low volatility organic compounds (ULVOC, C(300 K) < 3·10⁻⁹ μg m⁻³) were observed.”*

The following is in our discussion of contribution to growth

“Early stage particle growth is therefore plausibly driven by measured LVOC and ELVOC as these will be the molecules involatile enough to readily condense down on particles of this size. From 2D-VBS volatility calculations discussed in the previous section, we show that the LVOC and ELVOC we measure in Barcelona plausibly arise from the oxidation of aromatics (particularly PAHs in the case of ELVOC) and monoterpenes.”

And in our conclusions and abstract respectively, state

“SVOC arose from mostly isoprene and alkylbenzene oxidation, whereas LVOC and ELVOC arose from alkylbenzene, monoterpene and PAH oxidation”

“The concentration of these HOMs shows a dependence on temperature, and HOMs primarily fall into the SVOC volatility class. LVOC arise from oxidation of alkylbenzenes, PAHs and monoterpenes, whereas ELVOC appear to arise from primarily PAH and monoterpene oxidation”

From these data we also calculate growth rates as according to Nieminen et al., (2010) from H₂SO₄ condensation, from condensation of H₂SO₄, MSA and HIO₃, which were present at low concentration, and from condensation of organics by volatility class. Although we understand binning growth by volatility as such is a rather large oversimplification, we got good agreement between our measured GR₅₋₂₀ from condensation of SVOC, LVOC, ELVOC, MSA, HIO₃ and H₂SO₄, and GR_{1.9-5} from LVOC, ELVOC, MSA, HIO₃ and H₂SO₄. The details of these calculations are included in the methods section, and are as follows

“Growth rates from irreversible condensation of various vapours were calculated according to the method of Nieminen et al. (2010). At our measured relative humidity, sulphuric acid favours binding to 2 H₂O molecules (Kúrten et al., 2007). As amine concentrations are likely limited, we presume no mass from amines in the condensing species. H₂SO₄ was assigned a density of 1.8 g cm⁻³. For simplicity, the properties of MSA regarding density and hydration were presumed the same as H₂SO₄, and HIO₃ was presumed to have the same hydration as H₂SO₄, with a density of 4.98 g cm⁻³. The density of condensing organic vapours was assumed to be 1.5 g cm⁻³, and concentration-weighted mean mass (~276 g mol⁻¹ for LVOC) and atomic weighted diffusion volumes of organic compounds were used to calculate GRs.”

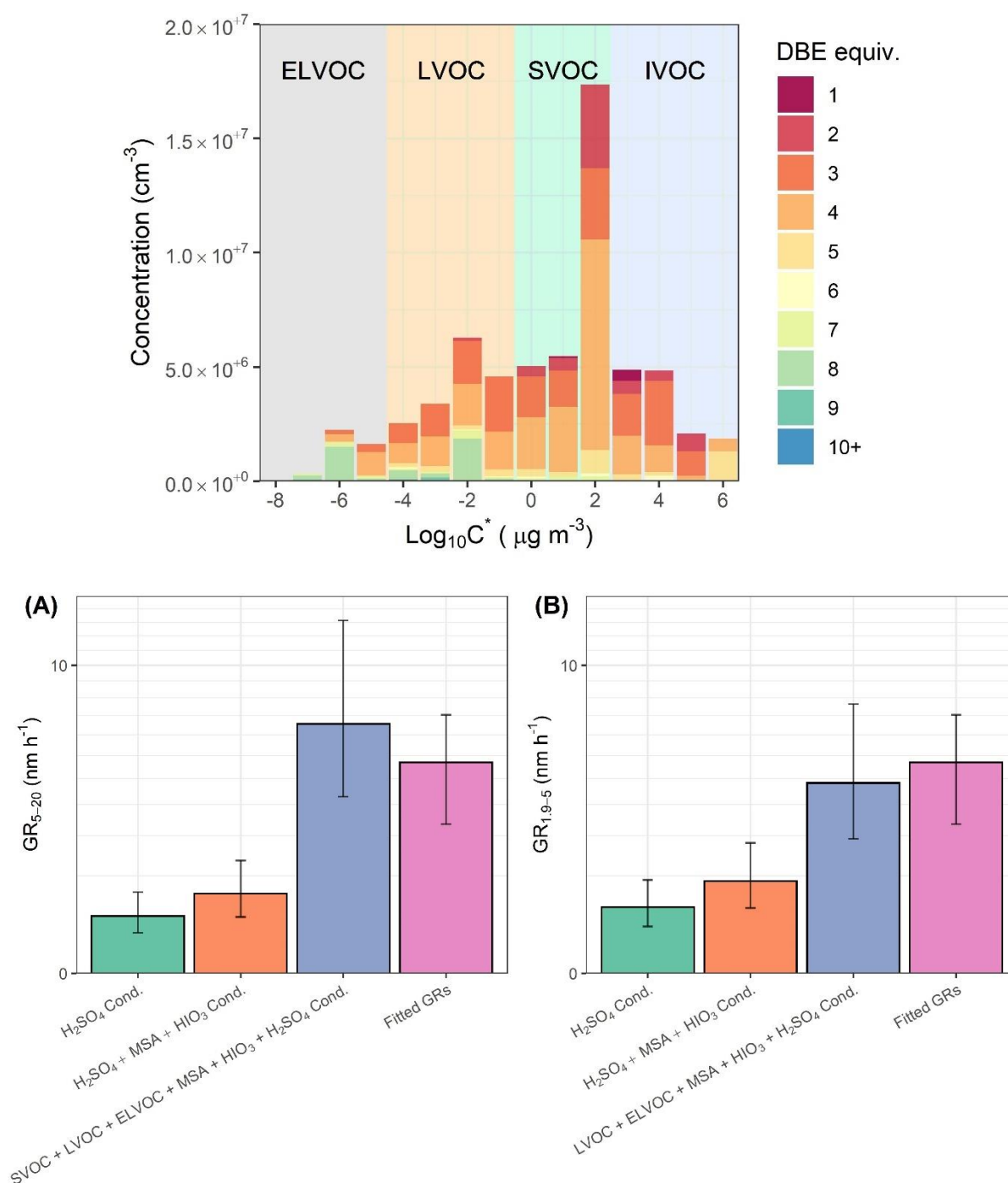
“The saturation vapour pressure at 300 K is defined by the 2D-volatility basis set (2D-VBS) as follows, if all nitrogen functionality is assumed to take the form -ONO₂ (Bianchi 2019; Donahue 2011; Schervish and Donahue, 2020):

$$\text{Log}_{10}(C^*)(300\text{ K}) = (N_{C_0} - N_C)b_C - N_O b_O - 2 \frac{N_O N_C}{N_C + N_O} b_{CO} - N_N b_N \quad (4)$$

Where N_C, N_H, and N_N, are the number of carbon, hydrogen, and nitrogen atoms respectively. N_O is the number of oxygen atoms minus 3N_N to account for -ONO₂ groups, N_{C0} is 25 (the carbon number of a 1 μg m⁻³ alkane), b_C, b_O, b_{CO}, and b_N are 0.475, 0.2, 0.9 and 2.5 respectively, and represent interaction and nonideality terms. The final term of equation (4) represents the -ONO₂ groups, each reducing the saturation vapour pressure by 2.5 orders of magnitude. C values are calculated at 300 K and not corrected, as 300 K is within 1 K of the campaign average temperature”*

We include the first of the below figures as a new Figure 8, and the growth-rate calculations as a supplementary figure. In light of these calculations we state the following in the text

“Calculated growth rates according to the method of Nieminen et al. (2010) are presented in Figure S6 for both GR_{1.9-5} and GR₅₋₂₀. Best agreement for GR₅₋₂₀ is when condensation of SVOC, LVOC, ELVOC, MSA, HIO₃ and H₂SO₄ is considered, and best agreement for GR_{1.9-5} is seen for condensation of all these apart from SVOC. The uncertainties in this method are large, and assumptions of irreversible condensation of SVOC onto particles of 5 nm likely lead to overestimations; however, these results confirm the essential role of the condensation of organic compounds to produce high growth rates observed in urban environments.”



Comment 5: The authors conclude: “We show new particle formation rates in Barcelona are linearly dependent upon the sulphuric acid concentrations, and this mechanism plausibly proceeds

by the formation of clusters involving sulphuric acid and highly oxygenated organic molecules, with likely involvement of bases”. Where do the authors see the clusters between organics and sulfuric acid in Figure 8? Figure 5 shows the H₂SO₄/DMA nucleation rates from CLOUD by Almeida et al. This data has been revised by Kürten et al. 2018. The new values would be at least an order of magnitude higher than Barcelona, which could eventually be explained by the higher ambient temperature in Barcelona.

Response 5: We unfortunately do not measure such clusters. This conclusion was based upon the observations of high formation rates matching those of Almeida et al. (2013), while simultaneously measuring a low sulphuric acid dimer:monomer ratio. The root cause of this, we hypothesised, could be the simultaneous involvement of sulphuric acid, bases and HOMs, however, in light of the updated Figure 6 utilising the data of Kürten et al. (2018), we revise our conclusions a little. We state the following in section 3.2

“Nucleation rates measured in Barcelona ($J_{1.9} 178 \pm 190 \text{ cm}^{-3} \text{ s}^{-1}$ at $[\text{H}_2\text{SO}_4] 7.1 \cdot 10^6 \pm 2.7 \cdot 10^6 \text{ cm}^{-3}$) are around an order of magnitude lower than that seen for the H₂SO₄-DMA-H₂SO₄ system, but exceed that of the H₂SO₄-BioOxOrg-H₂O system by ~1 order of magnitude, and that of the H₂SO₄-NH₃-H₂O and H₂SO₄-H₂O system by multiple orders of magnitude. No dissimilarity is seen between the data points corresponding to full or burst type nucleation, indicating similar mechanisms of formation, despite lower HOM concentrations on burst-event days”

“Particle formation plausibly operates by sulphuric acid-amine nucleation involving the measured C2 and C4 amines in our data, with nucleation rates hindered relative to those measured in the CLOUD experiments by elevated temperatures, and a decline to the sulphuric acid dimer:monomer ratio indicates that base concentrations may be limited. We cannot rule out an involvement of HOMs in particle formation processes, and, as no higher-order clusters were observed, we cannot establish sulphuric acid-amine nucleation with certainty.”

Comment 6: *From Figure 7a the authors claim a temperature dependence of HOM formation. However, the higher HOM concentrations at high temperature are also accompanied by higher global radiation. Thus, this dependence could just represent day-night time chemistry and the dependence on OH concentration.*

Response 6: If we take just the daylight hours, or hours where insolation is $>100 \text{ W m}^{-2}$, the correlation between temperature and signal gets slightly better $R^2 = 0.29 \rightarrow R^2 = 0.30$. Taking the periods where insolation is $<100 \text{ W m}^{-2}$ then there is no correlation seen. However, the correlation between HOMs and insolation is very poor ($R^2 = 0.15$). We maintain that there exists a dependence of HOM concentration on temperature within this dataset.

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4 **Molecular Insights into New Particle Formation**
5 **in Barcelona, Spain**
6

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27 ABSTRACT

28 Atmospheric aerosols contribute some of the greatest uncertainties to estimates of global radiative
29 forcing, and have significant effects on human health. New particle formation (NPF) is the process
30 by which new aerosols of sub-2 nm diameter form from gas-phase precursors and contributes
31 significantly to particle numbers in the atmosphere, accounting for approximately 50% of cloud
32 condensation nuclei globally. Here, we study summertime NPF in urban Barcelona in NE Spain
33 utilising particle counting instruments down to 1.9 nm and a Nitrate CI-API-ToF. The rate of
34 formation of new particles is seen to increase linearly with sulphuric acid concentration-, although
35 particle formation rates fall short of chamber studies of H₂SO₄-DMA-H₂O, while exceeding those
36 of H₂SO₄-BioOxOrg-H₂O nucleation, although a role of highly oxygenated molecules (HOMs)
37 cannot be ruled out in a manner similar to systems studied in chamber studies involving sulphuric
38 acid, water and dimethylamine (DMA), as well as sulphuric acid, water and the oxidation products
39 of pinanediol. The sulphuric acid dimer:monomer ratio is significantly lower than that seen in
40 experiments involving sulphuric acid and DMA in chambers, indicating that stabilization of
41 sulphuric acid clusters by bases is weaker in this dataset than in chambers, either due to rapid
42 evaporation due to high summertime temperatures, or limited pools of stabilising amines and thus
43 another mechanism. Such a mechanism cannot be verified in this data, as no higher-order H₂SO₄-
44 amine clusters, nor H₂SO₄-HOM clusters were measured, likely involving the plentiful highly
45 oxygenated organic molecules (HOMs) is plausible. The high concentrations of HOMs arise largely
46 from isoprene, both alkylbenzene, and monoterpene and PAH oxidation, with alkylbenzenes the
47 former providing greater concentrations of HOMs due to significant local sources. The
48 concentration of these HOMs shows a dependence on both temperature and precursor VOC
49 concentration. The organic compounds measured primarily fall into the SVOC volatility class
50 arising from alkylbenzene and isoprene oxidation. LVOC largely arise from oxidation of
51 alkylbenzenes, PAHs and monoterpenes, whereas ELVOC arise from primarily PAH and
52 monoterpene oxidation. New particle formation without growth past 10 nm is also observed, and on

53 these days ~~the highly oxygenated organic compound~~oxygenated organic concentrations ~~is are~~
54 ~~significantly~~ lower than on days with growth by a factor of 1.6, and thus high concentrations of low
55 volatility oxygenated organics which primarily derive from traffic-emitted VOCs appear to be a
56 necessary condition for the growth of newly formed particles in Barcelona. These results are
57 consistent with prior observations of new particle formation from sulphuric acid-amine reactions in
58 both chambers and the real atmosphere, and ~~these results~~ are likely representative of the urban
59 background of many European Mediterranean cities. A role for HOMs in the nucleation process
60 cannot be confirmed or ruled out, and there is strong circumstantial evidence for the participation of
61 higher molecular weight HOMs-HOMs across multiple volatility classes in particle growth.

62

63 1. INTRODUCTION

64 Atmospheric aerosols, defined as liquid or solid droplets suspended in a gas, affect the climate both
65 directly by scattering and absorbing radiation, and indirectly by acting as cloud condensation nuclei
66 (CCN) (Penner et al., 2011), providing great uncertainties in estimates of global radiative forcing
67 (IPCC, 2014). Further, fine ambient aerosols (defined as those with diameter below 2.5 μm) are the
68 fifth greatest global mortality risk factor, resulting in 103.1 million disability-adjusted life year loss
69 in 2015 (Cohen et al., 2017). The number concentration of the ultrafine fraction of these (aerosols
70 with diameter below 0.1 μm , referred to as ultrafine particles or UFP) pose potentially significant
71 health risks also, due to their high concentration and surface area. The more diffuse, gas-like
72 behaviour of UFP allows them to penetrate into the deep lung and enter the bloodstream (Miller et
73 al., 2017). Ultrafine particles occur in the urban environment either as primary emissions (~~i.e.e.g.~~,
74 from car exhaust (Harrison et al., 2018)) or secondarily as the product of new particle formation
75 (NPF) ([Brines et al., 2015](#); [Guo et al., 2014](#); [Kulmala et al., 2017](#); [Lee et al., 2019](#)) (~~Agudelo-~~
76 ~~Castañeda et al., 2019~~; ~~Bousiotis et al., 2019~~)
77
78 NPF is the formation of aerosol particles from gas-phase precursors. NPF can be considered a two-
79 step process involving initial formation of a cluster of gas molecules at the critical diameter at
80 around 1.5 nm - the diameter at which a free-energy barrier must be overcome to allow the
81 spontaneous phase transition from gas to liquid or solid (Zhang et al., 2012), and the subsequent
82 growth of this droplet to a larger aerosol particle. The first step of this process is dependent upon
83 the stability and abundance of the clustering molecules. Sulphuric acid, water, and dimethylamine
84 (DMA), for example, efficiently form particles as the strong hydrogen bonding between the acid
85 base pair produces near negligible evaporation, much lower than the evaporation rate seen for the
86 more weakly bound sulphuric acid-ammonia-water system. Nucleation of sulphuric acid, DMA and
87 water proceeds at, or near to the kinetic limit [in a chamber at 278 K when DMA mixing ratios are](#)
88 [sufficient](#) (Almeida et al., 2013; [Kürten et al., 2014](#)~~Kurtén et al., 2008~~). Once past this 1.5 nm

89 diameter, condensation and coagulation will drive particle growth. Both the abundance of
90 condensable gases and their vapour pressures limit condensational growth. Vapour pressures are
91 especially important for the initial growth stages, as the Kelvin effect barrier impairs condensation
92 of more volatile species, with this condition of low vapour pressures becoming less significant as
93 the diameter of the particle increases (Tröstl et al., 2016). Once sufficiently large (>50 nm), the loss
94 processes [of coagulation and evaporation](#) ~~for~~ of these particles become inefficient, resulting in a
95 significant atmospheric lifetime. It is ~~at~~ [from these](#) these diameters [onwards](#) the climate forcing
96 effects of these particles become most pronounced.

97

98 NPF processes happen globally, across a diverse range of environments from pristine polar regions,
99 to polluted urban megacities (Kerminen et al., 2018), and represent a significant source of CCN,
100 with 10-60% of NPF events shown to produce CCN and enhancement factors to CCN count ranging
101 from 0.5 – 1:1 (Lee et al., 2019 and references within). Strong NPF events are observed across a
102 range of urban environments, despite high condensation sinks $>10^{-2} \text{ s}^{-1}$ (Bousiotis et al., 2019; Yu et
103 al., 2016), and can act as a precursor to strong haze events (Guo et al., 2014). ~~Urban NPF occurs~~
104 ~~despite extremely high condensation sinks, an effect which has~~ [The occurrence of urban NPF has](#)
105 only been partially explained by growing understanding from recent in-depth studies (Yao et al.,
106 2018). Recent advances in instrumentation allow for the measurement of particles down to the
107 critical diameter with instruments such as the particle size magnifier (PSM), and (Neutral) Air Ion
108 Spectrometer (NAIS/AIS) (Lee et al., 2019), and mass spectral techniques for measuring the
109 abundance and composition of neutral (Jokinen et al., 2012) and charged (Junninen et al., 2010)
110 clusters. Elucidated mechanisms with these techniques involve sulphuric acid and ammonia in
111 remote environments (Jokinen et al., 2018; Yan, 2018), monoterpene derived highly oxygenated
112 molecules (HOM) in remote environments (Rose et al., 2018), iodic acid in coastal environments
113 (Sipilä et al., 2016), and sulphuric acid and DMA in polluted urban environments (Yao et al., 2018).

114

115 Urban Barcelona sees frequent, strong summer-time NPF events [occurring on 28% of days](#). These
116 events are associated with high insolation and [relatively elevated high-ozone \(\$\sim 60 \mu\text{g m}^{-3}\$ \)](#) when
117 considering the whole year (Brines et al., 2014, 2015). ~~This has been also reported in other urban~~
118 ~~environments such as Los Angeles, Brisbane, Rome and Madrid (Brines et al., 2015). However,~~
119 ~~when considering mid-summer, maximum ozone episodes are typically associated with high aerosol~~
120 ~~load, but not with NPF around the regional background of Barcelona (Camerero et al., 2019; Querol~~
121 ~~et al., 2017).~~ Ground-level observations report NPF events starting typically at midday, and either
122 occurring in urban Barcelona and the surrounding regional background simultaneously, or isolated
123 to ~~either just~~ urban Barcelona, or just the regional background (Dall'Osto et al., 2013). Vertical
124 profiles over urban Barcelona reveal that NPF occurs at higher altitudes, and starts earlier in the
125 day, as at a given altitude these events are not suppressed by early traffic peaks contributing to
126 particle load (Minguillón et al., 2015). Here, we examine [molecular level evidence for formation of](#)
127 [particles at the gas phase mass spectral evidence and particle formation rates at the](#) critical diameter
128 from sulphuric acid in Barcelona, with possible contribution from strong bases and highly
129 oxygenated organic molecules (HOMs), as well as factors influencing subsequent particle growth.

130

131 **2. METHODS**

132 **2.1 Sampling Site**

133 The Palau Reial site in Barcelona (41°23'15" N, 2°6'53.64" E) is representative of the urban
134 background of Barcelona, located at the Institute of Environmental Assessment and Water Research
135 (IDAEA-CSIC) in the north-west of the city. Sampling was performed from a container 20 m from
136 a low traffic road, and 200 m from the nearest main road (Avinguda Diagonal). Data were taken
137 from 2018/06/28 through 2018/07/18.

138

139 **2.2 Chemical Ionisation Atmospheric Pressure Interface Time of Flight Mass** 140 **Spectrometry**

141 The Aerodyne Nitrate Chemical Ionisation Atmospheric Pressure ~~Interface~~-interface Time of Flight
142 Mass Spectrometer (CI-APi-ToF) was used to make measurements of neutral oxygenated organic
143 compounds, organic and inorganic acids, bases, and their molecular clusters at high time resolution
144 with high resolving power. The ionization system charges molecules by adduct formation, such as
145 in the case of organic compounds with two or more hydrogen bond donor groups (Hytinen et al.,
146 2015), or proton transfer in the case of strong acids like sulphuric acid (Jokinen et al., 2012).
147 Hydroxyl or hydroperoxyl functionalities are both common hydrogen bond donating groups, with
148 hydroperoxyl being the more efficient hydrogen bond donor (Møller et al., 2017). This instrument
149 has been explained in great detail elsewhere (Jokinen et al., 2012; Junninen et al., 2010), but briefly,
150 the front end consists of a chemical ionisation system where a 10 L min⁻¹ sample flow is drawn in
151 through the 1 m length 1" OD stainless steel tubing opening. A secondary flow was run parallel and
152 concentric to this sample flow, rendering the reaction chamber effectively wall-less. A 3 cm³ min⁻¹
153 flow of a carrier gas (N₂) is passed over a reservoir of liquid HNO₃, entraining vapour which is
154 subsequently ionised to NO₃⁻ via an X-ray source. ~~This flow~~Ions is-are then guided into the sample
155 flow. The nitrate ions will then charge molecules either by clustering or proton transfer. The mixed
156 flows travelling at 10 L min⁻¹ enter the critical orifice at the front end of the instrument at 0.8 L min⁻¹
157 and are guided through a series of differentially pumped chambers before reaching the ToF
158 analyser. All data analysis was carried out in the Tofware package in Igor Pro 7 (Tofwerk AG,
159 Switzerland). Signals except for those of amines and ammonia are divided by the sum of reagent
160 ion signals and multiplied by a calibration coefficient to produce a concentration. Sensitivity-A
161 calibration coefficient of 3×10⁹ cm⁻³ was established based upon comparison with a sulphuric acid
162 proxy (Mikkonen et al., 2011) and is in line with a prior calibration with our instrument (Brean et
163 al., 2019). Uniform sensitivity between H₂SO₄ and all other species measured by CI-APi-ToF bar
164 amines and ammonia was assumed in this work. This introduces some uncertainties, as it relies
165 upon both collision rates and charging efficiencies to be the same within the ionisation source for all
166 species. Amine and ammonia signals are normalised to the nitrate trimer signal (Simon et al., 2016).

167 [Prior reports of ammonia and amines as measured by CI-APi-ToF employed corona discharge](#)
168 [systems, which utilise higher concentrations of nitric acid, thus we report normalised signals. We](#)
169 [present correlations of each of these bases clustered with the nitrate dimer plotted against](#)
170 [measurements with the nitrate trimer, as well as their intercorrelations and example peak fits across](#)
171 [Figure S1. C₂ amines, C₄ amines and ammonia were the only molecules of this kind found in our](#)
172 [mass spectra. assumed based upon a prior calibration \(Brean et al., 2019\) and comparison with a](#)
173 [calculated sulphuric acid proxy \(Mikkonen et al., 2011\). Systematic uncertainties of +100% / -50](#)
174 [% arising from this method are assumed.](#)

175 Due to the high resolving power of the CI-APi-ToF system (mass resolving power of 3000, and
176 mass accuracy of 20 ppm at 201 m/Q), multiple peaks can be fit at the same unit mass and their
177 molecular formulae assigned. Beyond 500 m/Q, peak fitting and assignment of compositions
178 becomes problematic as signal decreases, mass accuracy decreases, and the total number of possible
179 chemical compositions increases, so peaks above the C₂₀ region have not been assigned (Cubison
180 and Jimenez, 2015), however, signals past this region tended to be extremely low. [All ions](#)
181 [identified are listed in Table S1.](#) As proton transfer mostly happens with acids, and nearly all HOM
182 molecules will be charged by adduct formation it is possible to infer the uncharged formula;
183 therefore, all HOMs from here onwards will be listed as their uncharged form. The CI-APi-ToF
184 inlet was placed approximately 1.5 m a.g.l. CI-APi-ToF data is only available between the dates
185 2018/07/06 and 2018/07/17.

186

187 **2.2 Particle Size and Number Measurements**

188 Two Scanning Mobility Particle Sizer (SMPS) instruments measured particle size distributions at 5
189 minute time resolution, one Long Column SMPS (TSI 3080 EC, 3082 Long DMA, 3772 CPC, TSI,
190 USA) and one NanoSMPS (3082 EC, 3082 Nano DMA, 3776 CPC, TSI, USA) measuring the
191 ranges 10.9 – 478.3 nm and 4.5 – 65.3 nm respectively. A Particle Size Magnifier (A10, Airmodus,
192 FN) linked to a CPC (3775, TSI, USA) measured the sub-3 nm size fraction. The PSM was run in

193 stepping mode, operating at four different saturator flows to vary the lower size cut of particles that
194 it will grow (defined as the point of 50% efficiency, D_{50}). ~~This was set at 4 flows giving The~~
195 ~~instrument provided~~ D_{50} from 1.4 to 2.4 nm. The instrument switched between saturator flows each
196 2.5 minutes, giving a sub-2.4 nm size distribution every 10 minutes. Aerosol sampling inlets were
197 placed approximately 2 m a.g.l.

198

199 **2.3 Other Measurements**

200 Mixing ratios of non-methane VOC with proton affinity greater than H_3O^+ were made using the
201 proton transfer reaction time of flight mass spectrometer (PTR-ToF-MS 8000, Ionicon Analytik
202 GmbH, Austria). A detailed description of the instrument is provided by Graus et al., (2010) The
203 sampling set up, operating conditions, and quantification procedures are similar to those described
204 in Minguillón et al. (2016). Continual monitoring of composition and mass of submicron aerosol
205 ~~>75 nm~~ was carried out with an Aerosol Chemical Speciation Monitor (ACSM, Aerodyne, USA)
206 (Ng et al., 2011). Ozone, NO, and NO_2 were measured by conventional ultraviolet and
207 chemiluminescence air quality instrumentation. Meteorological data were supplied by the Faculty of
208 Physics of University of Barcelona, from a nearby (200 m from the measurement site)
209 meteorological station located at the roof of an 8 floor building.

210

211 **2.4 Condensation Sink and Particle Growth Rate**

212 The condensation sink (CS) represents the rate at which a vapour phase molecule will collide with
213 pre-existing particle surface, and was calculated from the size distribution data as follows ([Kulmala](#)
214 [et al., 2012](#)):

215

$$216 \quad CS = 4\pi D \sum_{d_p} \beta_{m,d_p} d_p N_{d_p}, \quad (1)$$

217

218

219 where D is the diffusion coefficient of the diffusing vapour (assumed sulphuric acid), β_m is a
220 transition regime correction (Kulmala et al., 2001), d_p is particle diameter, and N_{d_p} is the number of
221 particles at diameter d_p . The formation rate of new particles at size d_p is calculated as follows:

$$222 \quad J_{d_p} = \frac{dN_{d_p}}{dt} + CoagS_{d_p} \cdot N_{d_p} + \frac{GR}{\Delta d_p} \cdot N_{d_p} \quad (2)$$

224
225 where the first term on the right-hand side comprises the rate at which particles enter the size d_p ,
226 and the latter two terms represent losses from this size by coagulation and growth respectively. [J₅](#)
227 [was calculated using the data between 5 - 10 nm, and J_{1.9} was calculated using the measurements](#)
228 [between 1.9 – 4.5 nm. We also calculated J_{1.9} from our NanoSMPS data, employing the equations](#)
229 [of Lehtinen et al. \(2007\). J_{1.9} from both methods showed reasonable agreement \(\$R^2 = 0.34\$ \).](#)
230 [Agreement between J₅ and J_{1.9} for each method was similar \(\$R^2 = 0.37\$ and \$R^2 = 0.38\$ for J_{1.9}](#)
231 [calculated from PSM data and from Lehtinen et al. \(2019\) respectively\). J_{1.9} is greater than J₅ as](#)
232 [predicted from equation \(2\) by around a factor of 20. See Kulmala et al. \(2001\) for more](#)
233 [information on calculation of coagulation sinks and formation rates. Growth rates between 4.5 – 20](#)
234 [nm were calculated according to the lognormal distribution function method \(Kulmala et al., 2012\),](#)
235 [whereas those between 1.9 and 4.5 nm were calculated from PSM data using a time-delay method](#)
236 [between PSM and NanoSMPS data. Systematic uncertainties on our calculated J_{1.9} values include](#)
237 [25% method uncertainty \(Yli-Juuti et al., 2017\), with a further 25% arising from uncertainties in](#)
238 [PSM cutoff \(\$\pm 0.5\$ nm\), as well as a 10% uncertainty in counting errors. A 50% error arising from](#)
239 [calculated coagulation sink is also applied \(Kurten et al., 2016\). The above calculations rely on the](#)
240 [assumption of homogeneous air masses, and while air mass advection, as well as primary particle](#)
241 [emissions can cause errors in estimations of temporal changes in particle count and diameter, the](#)
242 [appearance and persistence of a new mode of particles across a period of several hours is typically](#)
243 [indicative of a regional process.](#)

245 Growth rates from irreversible condensation of various vapours were calculated as according to
 246 Nieminen et al. (2010). At our measured relative humidity, sulphuric acid favours binding to 2 H₂O
 247 molecules (Kürten et al., 2007). As amine concentrations are likely limited, we presume no mass
 248 from amines in the condensing species. H₂SO₄ was assigned a density of 1.8 g cm⁻³. For simplicity,
 249 the properties of MSA regarding density and hydration were presumed the same as H₂SO₄, and
 250 HIO₃ was presumed to have the same hydration as H₂SO₄, with a density of 4.98 g cm⁻³. The
 251 density of condensing organic vapours was assumed 1.5 g cm⁻³, and concentration-weighted mean
 252 mass (~276 g mol⁻¹ for LVOC) and atomic weighted diffusion volumes of organic compounds were
 253 used to calculate GRs.

254 and Kulmala et al. (2012) for calculation of growth rates.

255 2.4 DBE and 2D-VBS

256 The double bond equivalent (DBE) describes the degree of unsaturation of an organic molecule and
 257 is defined simply as:

$$259 \text{ DBE} = N_C - \frac{N_H}{2} - \frac{N_N}{2} + 1 \quad (3)$$

260 T

261 he saturation vapour pressure at 300 K is defined by the 2D-volatility basis set (2D-VBS) as
 262 follows, if all nitrogen functionality is assumed to take the form -ONO₂ (Bianchi 2019; Donahue
 263 2011; Schervish and Donahue, 2020):

$$265 \text{ Log}_{10}(C^*)(300 \text{ K}) = (N_{C0} - N_C)b_C - N_O b_O - 2 \frac{N_O N_C}{N_C + N_O} b_{CO} - N_n b_N \quad (4)$$

266

267 Where N_C, N_H, and N_N, are the number of carbon, hydrogen, and nitrogen atoms respectively. N_O is
 268 the number of oxygen atoms minus 3N_N to account for -ONO₂ groups, N_{C0} is 25 (the carbon
 269 number of a 1 μg m⁻³ alkane), b_C, b_O, b_{CO}, and b_N are 0.475, 0.2, 0.9 and 2.5 respectively, and
 270 represent interaction and nonideality terms. The final term of equation (4) represents the -ONO₂

271 [groups, each reducing the saturation vapour pressure by 2.5 orders of magnitude. \$C^*\$ values are](#)
272 [calculated at 300 K and not corrected for temperature, as 300 K is within 1 K of the campaign](#)
273 [average temperature.](#)

274

275 **3. RESULTS AND DISCUSSION**

276 **3.1 General Conditions of NPF Events**

277 Summer NPF events in the regional background around Barcelona are associated with high
278 insolation, relatively low ozone (high compared with the rest of the year), and lower particulate
279 matter load (Brines et al., 2014; Carnerero et al., 2019). Figure 1 shows an example of a day with no
280 NPF in panel (a), referred to as “non-event” here, where two traffic-associated peaks in particle
281 number are seen during rush hours. Midday traffic peaks are also seen on certain days, but these are
282 easily distinguished from nucleation processes due to the lack of a significant <10 nm mode. Figure
283 1(b) shows a nucleation day with growth to larger sizes >10 nm, termed “full-event”, showing the
284 growth through the course of the day. These fulfil all the criteria of Dal Maso et al. (2005). [4 events](#)
285 [of this type were observed with CI-APi-ToF data coverage.](#) Figure 1(c) shows a day with nucleation
286 occurring, but no growth past 10 nm. These days are referred to as “burst-event” days. Here, NPF is
287 seen to occur, but particles fail to grow past the nucleation mode. 2 such events were seen in this
288 data [with CI-APi-ToF data coverage](#), and both are accompanied by a distinct mode appearing
289 beforehand at ~20-40 nm. Condensation sinks were not significantly higher than on full event days,
290 so this failure of particles to grow further cannot be attributed to condensational [\(or coagulation\)](#)
291 losses. [\$GR_{4.5-20}\$ ranged between 2.47 – 7.31 \$\text{nm h}^{-1}\$ \(\$4.69 \pm 2.03 \text{ nm h}^{-1}\$ \), \$GR_{1.9-4.5}\$ ranged between](#)
292 [3.12 – 5.20 \$\text{nm h}^{-1}\$ \(\$4.36 \pm 1.02 \text{ nm h}^{-1}\$ \). The survival parameter \(P\) as suggested by Kulmala et al.](#)
293 [\(2017\) is defined as \$CS \cdot 10^{-4} / GR\$, and for this data is equal to 82, higher than other European cities.](#)
294 [The occurrence of such a high P value should, in theory, inhibit the occurrence of NPF, but we](#)
295 [show events happen readily under such conditions, akin to other heavily polluted megacities.](#)

296

297 Figure 2 contains box plots showing condensation sink, temperature and global radiation for all 3
298 NPF types [across the entire day \(diurnal profiles plotted in Figure S2\)](#). Condensation sinks during
299 NPF periods of both types (Figures 1(b) & 1(c)) were not significantly lower than in non-event
300 periods. [Condensation sinks were suppressed prior to the beginning of an event for full-events,](#)
301 [increasing relative to non-events through the afternoon period. Of the two burst-events, one was](#)
302 [similarly characterised by a suppression to condensation sink, whereas the other showed a sharp rise](#)
303 [in the midday](#). Global radiation and temperature were higher for full-events, most significantly for
304 temperature. Figure 3 is as Figure 2 but for sulphuric acid, ammonia and amines, and HOMs as
305 measured by CI-APi-ToF [\(HOM criteria are discussed in section 3.3.1\)](#). Sulphuric acid is elevated
306 during both full-event and burst-event periods. In urban Barcelona, sulphuric acid will primarily
307 arise from oxidation of SO₂ by the OH radical, with anthropogenic emissions such as shipping
308 emissions from the port areas being significant sources of SO₂ (Henschel et al., 2013). Direct traffic
309 emissions have been shown to be a significant primary sulphuric acid source (Olin et al., 2020), but
310 our sulphuric acid data show no traffic peaks. Ammonia and amines show enhancement for full-
311 event periods, but not burst-event periods. Nucleation rates (at typical tropospheric sulphuric acid
312 concentrations) are sensitive to amine concentrations in the range of a few pptv, with enhancements
313 to amine mixing ratios past this point increasing the nucleation rate marginally (Almeida et al.,
314 2013), [while typical](#) concentrations of DMA and other alkylamines vary from zero to a few pptv
315 in the boundary layer [\(Ge et al., 2011a\)](#).

316

317 Barcelona has been shown to contain ppbv levels of ammonia (Pandolfi et al., 2012), arising from
318 both agriculture to the north (Van Damme et al., 2018), and anthropogenic activities such as waste
319 management and traffic, with waste management being the primary ammonia source. Highest
320 ammonia mixing ratios are found in the densely populated old city centre (Reche et al., 2015).
321 Agriculture, waste management, and traffic are also all significant sources of low molecular weight
322 alkylamines, such as DMA (Ábalos et al., 1999; Cadle and Mulawa, 1980; Hutchinson et al., 1982;

323 Ge et al., 2011a), and are likely the source of amines found in this dataset. Activities such as
324 composting and food industry are especially strong sources of trimethylamine (TMA) (Ge et al.,
325 2011a). ~~A similar tertiary amine, trimethylamine (TEA) has been shown to be highly inefficient at
326 forming particles with sulphuric acid, likely due to its inability to form strong hydrogen bonded
327 pairs (Glasee et al., 2015), and NO₃⁻ ionization is therefore likely not sensitive to tertiary amines.
328 As the charging mechanism of amines is similar to the mechanism by which amines form clusters
329 with sulphuric acid, the sensitivities will be roughly similar to the strength of the bonding within
330 sulphuric acid clusters, therefore, although significant levels of TMA are expected, they were not
331 seen in the CI-APi-ToF spectra, nor will they participate efficiently in nucleation.~~ Although high
332 emission fluxes of TMA are expected in this environment, they are not present in our spectra. The
333 TMA ion has been reported ~~prior~~ previously with a similar ionisation setup to that utilised in this
334 study (Kürten et al., 2016) ~~The quantification of amines seen in the CI-APi-ToF is highly uncertain,
335 with sensitivities likely being dependent upon the efficiency with which an amine will form a
336 cluster with the nitrate trimer in the CI-APi-ToF, so no mixing ratios have been calculated from
337 these signals (Simon et al., 2016). All ammonia and amine signals follow similar trends, indicating
338 similar sources (R² ranging from 0.58—0.94 between ammonia and amines). On full-event days,
339 the signal for C₂ and C₄ amines has a midday elevation concurrent with peaks to solar radiation
340 (Figure S2), and can help explain the high formation rates we see in this dataset (see section 3.2).~~
341 The relative strength of these signals are shown in Figure S3,4, with significantly higher signals
342 attributed to ammonia compared to amines, despite a likely lower sensitivity (Simon et al., 2016).
343
344 HOM concentrations were greatly enhanced during full-event periods (factor of 1.5 higher
345 compared to non-NPF mean), but lower during burst-event periods (factor of 1.2 lower compared to
346 non-NPF mean), implying their necessity for growth. The sources and implications of these HOMs
347 are discussed in section 3.3h. ~~Further, all full and burst events are associated with southerly and
348 south westerly air masses. Despite this,~~ concentrations of iodine and DMS-derived acids such as

349 iodic acid (HIO₃) and methanesulphonic acid (MSA) are low (7.8·10⁵ and 3.3·10⁵ cm⁻³
350 respectively), indicating a small influence of oceanic emissions on particle nucleation/growth.
351 Extended box plots as Figures 2 & 3 are presented in Figure S4, and HYSPLIT back trajectories per
352 event in Figure S5.

353

354 **3.2 Mechanisms of New Particle Formation**

355 The correlation between J_{1.95} and ~~ion signals for concentration of~~ sulphuric acid is plotted in Figure
356 4. ~~J₅ is used here in place of J_{1.5} due to better data coverage.~~ A close relationship between **NPF**
357 nucleation rates and sulphuric acid concentrations ($R^2 = 0.49$) are consistent with observations
358 globally (Lee et al., 2019). This relationship is not dependent upon condensation sink. These NPF
359 rates have no dependence on other ions as measured by CI-API-ToF, including HIO₃, MSA,
360 ammonia, amines or HOMs (R^2 for all <0.1). This is not to say that all of these molecules are not
361 involved in the nucleation process, rather that elevations or reductions to their concentrations during
362 nucleation periods do not have significant impact on NPF-nucleation rates. In the example of
363 alkylamines, their gas phase concentration may decrease due to clustering with elevated sulphuric
364 acid, as they cluster at around a 1:1 ratio at high amine mixing ratios (Kürten et al., 2014) (and
365 therefore they will not be detectable as free amines). Further, if amines are present at a few pptv,
366 their mixing ratios are significantly higher than our ambient measured sulphuric acid
367 concentrations, and will be sufficient to accelerate nucleation rates (Almeida et al., 2013). ~~and~~
368 ~~losses due to enhanced photochemistry~~ Photochemical losses will also be greater during the periods
369 of highest NPF rate (Ge et al., 2011b). ~~Concentrations for these ions are shown in Figure S5.~~ The
370 strength of the relationship between sulphuric acid and nucleation rate has been quantitatively
371 reproduced in chamber studies involving the H₂SO₄-H₂O-DMA, and H₂SO₄-H₂O-BioOxOrg
372 system, accurately reproducing tropospheric observations of nucleation rates (Almeida et al., 2013;
373 Riccobono et al., 2014), although a later revision of the former shows nucleation rates at 278 K
374 exceeding typical tropospheric observations in the presence of high mixing ratios of DMA (Kürten

375 [et al., 2018](#)). A comparison between our data and results from the CLOUD chamber is presented in
376 Figure 5; included are the H₂SO₄-H₂O, H₂SO₄-NH₃-H₂O (Kirkby et al., 2011), H₂SO₄-H₂O-DMA
377 ([Kürten et al., 2018](#)[Almeida et al., 2013](#)) and H₂SO₄-BioOxOrg-H₂O systems (Riccobono et al.,
378 2014) – BioOxOrg refers to the oxidation products of pinanediol (C₁₀H₁₈O₂) and OH. [Data from](#)
379 [these chamber experiments is for 278 K and 38 - 39 % relative humidity. The relationship between](#)
380 [sulphuric acid and nucleation rate is similar to the H₂SO₄-DMA-H₂SO₄ system, Nucleation rates](#)
381 [measured in Barcelona \(\$J_{1.9} 178 \pm 190 \text{ cm}^{-3} \text{ s}^{-1}\$ at \$\[\text{H}_2\text{SO}_4\] 7.1 \cdot 10^6 \pm 2.7 \cdot 10^6 \text{ cm}^{-3}\$ \) are around an](#)
382 [order of magnitude lower than that seen for the H₂SO₄-DMA-H₂SO₄ system with extremely high](#)
383 [nucleation rates, but exceed that of the H₂SO₄-BioOxOrg-H₂O system by ~1 order of magnitude,](#)
384 [and that of the H₂SO₄-NH₃-H₂O and H₂SO₄-H₂O system multiple orders of magnitude. compared](#)
385 [to the other systems. The nucleation rate is also broadly similar also to the system involving](#)
386 [BioOxOrg.](#) No dissimilarity is seen between the data points corresponding to full or burst type
387 nucleation, indicating similar mechanisms of formation, despite lower HOM concentrations on
388 [these burst-event days. Conversely, research in remote boreal environments show that the](#)
389 [mechanism of nucleation can modulate dependent upon the H₂SO₄:HOM ratio \(Yan et al., 2018\).](#)
390 [Model studies of sulphuric acid-amine nucleation show a decline in nucleation rate with](#)
391 [temperature \(Almeida et al., 2013; Olenius et al., 2017\), as the evaporation rate of sulphuric acid-](#)
392 [amine clusters will increase with temperature \(Paasonen et al., 2012\). Conversely, evaporation rates](#)
393 [of such small clusters, and resultant nucleation rates tend to increase modestly with increases in](#)
394 [relative humidity, most pronounced at lower amine concentrations \(Almeida et al., 2013; Paasonen](#)
395 [et al., 2012\). Despite this, high nucleation rates at temperatures nearing 300 K have been reported](#)
396 [prior previously \(Kuang et al., 2008; Kürten et al., 2016\), although these tend to show a temperature](#)
397 [dependence \(Yu et al., 2016\). \(Paasonen et al., 2012\)\(Almeida et al., 2013; Olenius et al., 2017\). No](#)
398 higher-order sulphuric acid clusters, sulphuric acid-base clusters, nor sulphuric acid-HOM clusters
399 were visible in the mass spectral data, likely due to these being below the limit of detection of the
400 instrument (Jokinen et al., 2012), so cluster identity cannot be directly identified. Sulfuric acid

401 trimer stabilisation is dependent upon base abundance (Ortega et al 2012), and conversely,
402 sensitivity of nitrate CI-APi-ToF to sulfuric acid-base clusters is reduced due to the high base
403 content of such clusters (Jen et al., 2016).

404

405 To further explore the relationship between sulphuric acid clusters and the rate of nucleation, the
406 sulphuric acid dimer:monomer ratio is plotted in Figure 6. ~~Sulphuric acid dimer roughly represents~~
407 ~~the strength of sulphuric acid clustering in the nitrate CI-APi-ToF, as significant fragmentation of~~
408 ~~sulphuric acid base clusters occurs upon charging, with-~~ The sulphuric acid dimer:monomer ratio is
409 elevated by the presence of gas-phase bases such as DMA, and this elevation is dependent upon
410 both the abundances and proton affinities of such bases (Olenius et al., 2017). Upon charging,
411 evaporation of water and bases from sulphuric acid clusters occurs, and thus these are detected as
412 sulphuric acid dimer ~~the system~~ (Ortega et al., 2012, 2014). ~~Stronger bases result in more sulphuric~~
413 ~~acid dimer making it through the system, and thus this ratio is higher for systems involving bases~~
414 ~~that form stronger clusters with sulphuric acid.~~ The binding energy of the bisulphate-H₂SO₄ ion is
415 in excess of 40 kcal mol⁻¹ (Curtius et al., 2001), and thus minimal declustering of the dimer is
416 expected within the CI-APi-ToF instrument – however, declustering of higher order sulphuric acid
417 clusters has been shown to be sensitive to voltage tune (Passanati et al., 2019), and this likely
418 extends to the dimer also, and as such discrepancies between sets of results due to instrument setup
419 cannot be ruled out. The ratio of sulphuric acid dimer:monomer is also highly sensitive to
420 condensation sinks, with a difference in dimer concentration of approximately a factor of 4
421 expected at 10⁷ cm⁻³ between 0.001 s⁻¹ (a clean environment) and 0.03 s⁻¹ (condensation sinks
422 during these NPF events measured in this dataset) (Yao et al., 2018) and thus our low
423 dimer:monomer ratio can, in part, be explained by elevated condensation sinks. The dashed line
424 represents the ~~lower limit of ratio sulphuric acid~~ that would be seen due to ion induced clustering
425 (IIC) in the nitrate chemical ionisation system for a 50 ms reaction time (Zhao et al., 2010). The
426 sulphuric acid dimer:monomer ratio seen in the CLOUD H₂SO₄-DMA-H₂O system is plotted,

427 alongside our own data from Barcelona. The ratio from our own data is seen to be much lower than
428 that for the system purely involving DMA as a ternary stabilising species. Similarly, this ratio is
429 lower than for reports of H₂SO₄-DMA-H₂O nucleation in Shanghai (Yao et al., 2018), but is
430 markedly similar to reports in central rural Germany (Kürten et al., 2016). Similar to central
431 Germany, this ratio increases at lower sulphuric acid concentrations to a ratio more similar to the
432 H₂SO₄-DMA-H₂O system. A possible explanation for this is that at higher sulphuric acid
433 concentrations, the concentrations of stronger stabilising bases are insufficient to stabilise all
434 present sulphuric acid, with the higher end of the sulphuric acid concentrations seen in this data
435 roughly equivalent to 1 pptv sulphuric acid ($3 \times 10^7 \text{ cm}^{-3} = 1.2 \text{ pptv sulphuric acid}$). We also cannot
436 account for clustering due to naturally charged sulphuric acid in the atmosphere, but ion
437 concentrations in urban environments tend to be small due to efficient sink processes (Hirsikko et
438 al., 2011). Particle formation plausibly operates by sulphuric acid-amine nucleation involving the
439 measured C₂ and C₄ amines in our data, with nucleation rates hindered relative to those measured in
440 the CLOUD experiments by elevated temperatures, and a decline to the sulphuric acid
441 dimer:monomer ratio indicates that base concentrations may be limited. We cannot rule out an
442 involvement of HOMs in particle formation processes, and~~Particle formation therefore plausibly~~
443 ~~operates by mechanisms similar to the H₂SO₄-DMA-H₂O and H₂SO₄-BioOxOrg-H₂O systems~~
444 ~~which may occur in parallel, as no higher-order clusters were observed, we cannot establish~~
445 sulphuric acid-amine nucleation with certainty.

446

447 **3.3 HOMs and Growth**

448 3.3.1 HOM composition and sources

449 Barcelona, as a densely populated urban agglomerate, is distinct from ~~is characterised by higher~~
450 ~~NO_x than~~ the remote conditions under which HOMs have primarily been studied (Bianchi et al.,
451 2016, 2017; Schobesberger et al., 2013; Yan et al., 2016), and is characterised by elevated
452 temperatures, insolation and NO_x mixing ratios, as well as a diverse host of potential precursor

453 VOC. The first of these affects HOM yields significantly, as yields are highly dependent upon
454 temperature (Quéléver et al., 2019; Stolzenburg et al., 2018). Lower temperatures result in slower
455 H-abstractions, which will result in the likelihood of an RO₂· to undergo a different reaction
456 pathway, such as termination with HO₂· to increase (Praske et al., 2018). This is particularly
457 important in this study if there is a large energy barrier for the first or second H-abstraction taking
458 place, as this will determine the number of hydrogen bond donating groups, and therefore whether
459 the NO₃⁻ CI-APi-ToF is sensitive to a molecule or not. Elevated insolation will result in enhanced
460 photochemistry, and thus more rapid RO₂· formation rates, whereas elevated NO_x will produce more
461 HOM with nitrate ester functionality (Garmash et al., 2019; Rissanen, 2018), which tend towards
462 higher volatilities, and less efficient participation in particle formation (Ehn et al., 2014; Lehtipalo
463 et al., 2018), and growth (Yao et al., 2020).

464

465 Oxygenated volatile organic compounds (OVOC) are defined as species visible in the nitrate CI-APi-
466 ToF that do not classify as HOM. Here, the first of the three criteria provided by Bianchi et al. (2019),
467 that HOM must be formed by peroxy radical autoxidation, cannot be applied to define HOM, as
468 knowledge as to whether a molecule is a result of autoxidation requires sound knowledge of the
469 structure of the precursors, oxidants and peroxy radical terminators present, however, the number of
470 molecules observed with $nN = 2$ is around an order of magnitude lower than that for $nN = 1$, where
471 the primary source of multiple nitrogen functionalities would be multiple peroxy radical termination
472 reactions from NO_x, and therefore while multiple generations of oxidation have been shown to occur
473 in aromatics (Garmash et al., 2020), it is a small contributor to the concentration of what is classed as
474 HOM here. The second criterion to define HOM are that they must be formed in the gas phase under
475 atmospherically relevant conditions, which we deem appropriately fulfilled as all CI-APi-ToF
476 measurements are of gas phase compounds, and the final criterion is that HOM must contain more
477 than 6 oxygen atoms. To attempt to satisfy these criteria as best possible, the criteria of both
478 containing 6 oxygen atoms and 5 carbon atoms or greater and having an O:C ratio >0.6 is applied.

479 The diversified range of HOM precursors in Barcelona will be primarily anthropogenic in origin.
480 Averaged PTR-MS mixing ratios of different VOCs are presented in Figure S6. Figure 7(a)(a)
481 shows temperature plotted against the signal of HOM concentration plotted against temperature,
482 showing a dependence of HOM concentrations on temperature, with a lesser dependence on global
483 radiations. The precursors for these HOMs are presumed to be largely isoprene, alkylbenzenes, and
484 monoterpenes, and PAHs. The mean peak intensities assigned to alkylbenzene derived HOMs are
485 approximately a factor of two higher than those assigned to isoprene and monoterpene oxidation
486 across this entire campaign. In this data these VOC mixing ratios are, with the exception of
487 isoprene, not largely temperature dependent, with many of these HOMs forming under negligible or
488 zero insolation, and therefore very low OH[·] concentrations. These nighttime HOMs will not be
489 derived from the oxidation of aromatics, however, as rates of oxidation of alkylbenzenes by O₃ and
490 NO₃[·] are negligible (Molteni et al, 2018). These nighttime HOMs will therefore mostly be derived
491 from biogenic emissions which undergo more rapid nocturnal oxidation, and are likely transported
492 from inland by the land breeze during night (Millán, 2014; Querol et al., 2017).

493
494 Operating under the assumption that C₅, C₆, C₇, C₈, and C₉ HOMs primarily arise from isoprene,
495 benzene toluene, C₂-alkylbenzene C₃-alkylbenzene oxidation respectively (Massoli et al., 2018;
496 Molteni et al., 2018; Wang et al., 2017), HOM signals plotted against parent VOC concentration
497 indicate their dependence upon that VOC. Here, a C₇ HOM is thought to follow the formula C₇H₈
498 ₁₂O₅₋₁₀N₀₋₂. We have plotted HOM concentrations against VOC concentrations in Figure 7(b). C₁₀
499 HOMs are not included in these analyses as these may primarily arise from C₁₀H₁₂₋₁₄ alkylbenzene,
500 or monoterpene oxidation. HOM concentration appears mostly independent of VOC concentration,
501 with the exception of isoprene, for which emissions are highly temperature dependent, and thus this
502 is likely a function of the effect of temperature on HOM formation (Figure 7(a)). A lack of
503 correlation between other VOCs and their HOMs confirms that this relationship between HOMs
504 and temperature is not a function of enhanced VOC emission fluxes from, for example, evaporation,

505 except in the instance of isoprene. Fragmented monoterpene oxidation products will also contribute
506 to the total number of C₉ HOMs, and similarly, other VOCs can fragment upon oxidation. However,
507 these results indicate that HOM concentrations are elevated by temperature, and operate quite
508 independent of precursor VOC concentration.

509
510 DBE as calculated by equation 3 is equal to the number of pi bonds and rings within a molecule.
511 Benzene, toluene, and similar aromatics have DBE = 4, naphthalene = 7 and monoterpenes = 3.
512 DBE can be used as an indicator of sources when considering HOM in bulk. Saturation mass
513 concentration as calculated by equation 4 can help describe capacity of a molecule to both condense
514 onto newly formed particles and participate in nucleation. Figure 8 shows concentrations of HOMs
515 and other oxygenated organic molecules binned to the nearest integer Log₁₀(C*)(300 K), coloured
516 by DBE. Mean ion signals per carbon number are shown in Figure S7. Most measured molecules
517 fall into the SVOC class ($0.3 < C^*(300\text{ K}) < 300\ \mu\text{g m}^{-3}$) which will mostly exist in equilibrium
518 between gas and particle phase. Highest SVOC concentrations arise from fingerprint molecules for
519 isoprene oxidation under high NO_x (C₅H₁₀N₂O₈) (Brean et al., 2019), and oxidation of small
520 alkylbenzenes (C₇H₈O₅, C₈H₁₀O₅). LVOC and ELVOC ($3 \cdot 10^{-5} < C^* < 0.3\ \mu\text{g m}^{-3}$ and $3 \cdot 10^{-9} <$
521 C*(300 K) < $3 \cdot 10^{-5}\ \mu\text{g m}^{-3}$ respectively) have a greater contribution from molecules with higher
522 DBE, i.e., C₁₀H₁₀O₈ arising most likely from PAH oxidation (Molteni et al., 2018), and C₁₀H₁₅O₇N,
523 a common molecule arising from monoterpene oxidation in the presence of NO_x. The contribution
524 of molecules with carbon number ≤ 9 to these LVOC is modest, and ELVOCs are entirely
525 comprised of molecules with carbon numbers ≥ 10, and is dominated by DBEs of 8 and 4,
526 attributable to PAH and monoterpene oxidation respectively. No molecules classed as ultra-low
527 volatility organic compounds (ULVOC, C*(300 K) < $3 \cdot 10^{-9}\ \mu\text{g m}^{-3}$) were observed in our data, and
528 thus any pure HOM nucleation is unlikely. High insolation will also result in higher HO₂[•], as well as
529 higher RO₂[•] and other radicals. NO_x, HO₂[•] act as peroxy radical terminators, reducing the likelihood
530 of autoxidation to produce high O:C molecules, and the likelihood of RO₂[•]-RO₂[•] dimerization. The

531 $\text{NO}_x\text{-RO}_2^-$ reaction also produces HOMs with nitrate ester functionality (Brean et al., 2019;
532 Garmash et al., 2019; Rissanen, 2018), which tend to have higher volatilities, and are less efficient
533 at participating in early stage NPF (Ehn et al., 2014; Lehtipalo et al., 2018), likely due to
534 intramolecular H bonding (Elm et al., 2017). This, combined with plentiful VOCs with carbon
535 number <10 results in more volatile HOMs (see Figures S2 and S3) likely classed as LVOC or
536 SVOC (Bianchi et al., 2019; Tröstl et al., 2016). HOMs of this volatility are mostly incapable of
537 producing particles in the absence of sulphuric acid by ion induced mechanisms observed in a
538 chamber study and the remote environment (Kirkby et al., 2016; Rose et al., 2018). Peaks due to
539 certain night time HOMs are seen, largely C_{10} -HOMs attributable to inland air masses (Querol et
540 al., 2017). Many of these night time C_{10} -HOMs contain nitrogen functionalities, attributable to
541 either oxidation by NO_3^- radicals or NO_x chemistry and were not associated with night-time NPF.

542

543 As shown in Figure 3, an elevated HOM concentration appears to be a necessary condition for
544 particle growth past 10 nm during NPF events. These days are associated with elevated
545 temperatures, solar radiation, higher ozone, and lower $\text{NO}:\text{NO}_2$ ratio. HIO_3 is also significantly
546 higher on burst event days, likely a function of air mass trajectory, as HIO_3 would not inhibit
547 particle growth, condensational growth being a reversible, step-wise kinetic process. A recent study
548 in a remote environment reports growth rates matching condensation rates without accounting for
549 aqueous phase chemistry (Mohr et al., 2019). HOM yields are highly dependent upon temperature
550 (Quéléver et al., 2019; Stolzenburg et al., 2018). Lower temperatures result in slower H-
551 abstractions, which will result in the likelihood of an RO_2^- to undergo a different reaction pathway,
552 such as termination with HO_2^- to increase (Praske et al., 2018). This is particularly important if there
553 is a large energy barrier for the first or second H abstraction taking place, as this will determine the
554 number of hydrogen bond donating groups, and therefore whether the NO_3^- -CI-API-ToF is sensitive
555 to a molecule or not. Figure 7(a) shows temperature plotted against the signal of HOMs. The
556 precursors for these HOMs are presumed to be largely isoprene, alkylbenzenes and monoterpenes.

557 ~~The mean peak intensities assigned to alkylbenzene derived HOMs are approximately a factor of~~
558 ~~two higher than those assigned to isoprene and monoterpene oxidation across this entire campaign.~~
559 ~~In this data these VOCs are, with the exception of isoprene, not largely temperature dependent, with~~
560 ~~many of these HOMs forming under negligible or zero insolation, and therefore very low OH⁻~~
561 ~~concentrations. These nighttime HOMs will not be derived from the oxidation of aromatics,~~
562 ~~however, as rates of oxidation of alkylbenzenes by O₃ and NO₃⁻ are negligible (Molteni et al., 2018).~~
563 ~~These nighttime HOMs will therefore mostly be derived from biogenic emissions which undergo~~
564 ~~more rapid nocturnal oxidation, and are likely transported inland by the land breeze during night~~
565 ~~(Millán, 2014; Querol et al., 2017).~~

566
567 Operating under the assumption that C₅, C₇, and C₉ HOMs primarily arise from isoprene, toluene
568 and C₃ alkylbenzene oxidation respectively (Molteni et al., 2018; Wang et al., 2017), HOM signals
569 plotted against parent VOC concentration indicate their dependence upon that VOC. Here, a C₇
570 HOM is thought to follow the formula C₇H₈₋₁₂O₅₋₁₀N₀₋₂. This has been done in Figure 7(b). HOM
571 concentration is broadly dependent on VOC concentration. This is most significant for isoprene.
572 Fragmented monoterpene oxidation products will also contribute to the total number of C₉ HOMs,
573 and similarly, other VOCs can fragment upon oxidation. However, these results indicate that the
574 limiting factor in HOM production is temperature, and to a lesser degree VOC concentration.

575

576 3.3.2 HOMs and NPF

577 As shown in Figure 3, an elevated HOM concentration appears to be a necessary condition for particle
578 growth past 10 nm during NPF events. These days are associated with elevated temperatures, solar
579 radiation, higher ozone, and lower NO:NO₂ ratio. HIO₃ is also significantly higher on burst-event
580 days. A recent study in a remote environment reports growth rates matching condensation rates
581 without accounting for aqueous phase chemistry (Mohr et al., 2019). From 2D-VBS volatility
582 calculations discussed in the previous section, is it shown that LVOC and ELVOC measured in
583 Barcelona plausibly arise from the oxidation of aromatics (particularly PAHs in the case of ELVOC)

584 and monoterpenes. Calculated growth rates according to the method of Nieminen et al. (2010) are
585 presented in Figure S8 for both GR_{1.9-5} and GR₅₋₂₀. Best agreement for GR₅₋₂₀ is when condensation
586 of SVOC, LVOC, ELVOC, MSA, HIO₃ and H₂SO₄ is considered, and best agreement for GR_{1.9-5} is
587 seen for condensation of all these except SVOC. The uncertainties in this method are large, and
588 assumptions of irreversible condensation of SVOC onto particles of 5 nm likely lead to
589 overestimations; however, these results indicate an essential role of the condensation of organic
590 compounds to produce high growth rates observed in urban environments.

591

592 Figure 8-9 shows three mass-defect plots for a non-event period, full-event period, and burst-event
593 period (~~event days are the same days as shown in Figure 1~~). Here, oxygenated volatile organic
594 compounds (OVOC) are defined as species visible in the nitrate CI APi ToF that do not classify as
595 HOM. Here, the detailed criteria provided by Bianchi et al. (2019) cannot be applied to define
596 HOM, as knowledge as to whether a molecule is a result of autoxidation requires sound knowledge
597 of the structure of the precursors present. The criteria of both containing 6 oxygen atoms and 5
598 carbon atoms or greater, and having an O:C ratio >0.6 is applied, as these molecules will all
599 plausibly fulfil the updated criteria of "HOM". ~~The particular non-event~~ non-event day included in
600 Figure 8 was characterised by lower solar radiation and temperatures than average, so lower signals
601 for oxygenated species are seen due to weaker photochemistry (i.e., OH[•] concentration), and slower
602 autoxidation due to slower H-shift reactions (Frege et al., 2018; Quéléver et al., 2019; Stolzenburg
603 et al., 2018). The full-event day sees enhancements to smaller OVOCs and HOMs compared to the
604 non-event day, especially around 150-200 m/Q, which contains peaks corresponding to dicarboxylic
605 acids and isoprene oxidation products. Some of the largest peaks in the mass spectra correspond to
606 formulae seen arising from the enhanced OH[•] oxidation of alkylbenzenes (such as C₇H₇NO₆)
607 (Molteni et al., 2018; Wang et al., 2017). Larger HOMs see a less significant enhancement to
608 smaller alkylbenzene derived HOMs. ~~Most significant is t~~ The presence of many larger, unidentified
609 HOMs >4500 m/Q is enhanced during full-events, these peaks will comprise the largest

610 compounds, most likely of class ELVOC, arising from the oxidation of large VOCs, or RO₂-RO₂
611 accretion reactions, and thus, we likely underpredict ELVOC concentrations and resultant impacts
612 on particle growth in Figure S8. These unidentified peaks >400 m/Q are . During full-event periods,
613 ~~these peaks are~~ both more numerous and larger during full-event periods, with a factor of two
614 difference in total peak area. This area of the mass spectrum will compromise the largest HOMs,
615 from oxidation of large VOCs or from the RO₂ + RO₂ dimerization reaction of two smaller RO₂,
616 ~~these will likely have extremely low vapour pressures and be able to drive early stage particle~~
617 growth (Mohr et al., 2019; Tröstl et al., 2016). The burst-event day has significantly lower
618 concentrations of OVOCs and HOMs, and to a lesser degree, their nitrogen containing counterparts
619 (N-OVOCs and N-HOMs), with significantly fewer compounds >~~500~~400 m/Q. The most
620 significant difference between full and burst-event days is in the SVOCs, accounting for a factor of
621 two difference in concentration~~the HOMs on both types of event days is the C₉ and C₁₀ HOMs,~~
622 ~~consistent with lower concentrations of monoterpenes on burst-event days.~~ The sulphur containing
623 acids all have similar peak areas to the full-event day. These elevations to condensable OVOCs and
624 HOMs on particle formation days with growth are consistent with particle composition data as
625 measured by ACSM (Figure S9), ~~also showing a significant increase to organic mass concentration~~
626 ~~in the late evening (around when new particles reach sizes measurable by ACSM) Particle~~
627 composition on full-event days shows an elevation to organic mass concentration in the late evening
628 and night around when new particles from NPF will reach sizes detectable by the ACSM (~75 nm,
629 Ng et al., 2011). Organic mass between 16:00 – 23:00 is 3.5 µg m⁻³ on burst-event days, versus 7.8
630 µg m⁻³ on full-event days. and night on full-event days. This was not seen on days without
631 ~~nucleation, nor on days with nucleation but no growth (see Figure S4).~~

632

633 **4. CONCLUSIONS**

634 We show new particle formation rates in Barcelona are linearly dependent upon the sulphuric acid
635 concentrations, and while formation rates far exceed that of H₂SO₄-BioOxOrg-H₂O nucleation, they

636 fall short of those of H₂SO₄-DMA-H₂O nucleation at 278 K, as does the sulphuric acid
637 dimer:monomer ratio, possibly explained by cluster evaporation due to high temperatures in
638 summertime Barcelona (303 K during events), and limited pools of gas-phase amines. These results
639 are similar to reports of nucleation rates in rural Germany (Kürten et al., 2016). As no higher-order
640 clusters were directly measured, we cannot determine nucleation mechanisms with certainty, and an
641 involvement of HOMs in nucleation is plausible.

642
643 High concentrations of OVOCs and HOMs were measured by CI-APi-ToF. Of these, the ~~and this~~
644 mechanism plausibly proceeds by the formation of clusters involving sulphuric acid and highly
645 oxygenated organic molecules, with likely involvement of bases. This multiplicity of mechanisms
646 has been shown to occur in chamber studies but has not been observed in the real atmosphere
647 previously. The rate of nucleation relative to sulphuric acid concentration is similar to observations
648 of H₂SO₄-DMA-H₂O and H₂SO₄-BioOxOrg-H₂O nucleation. Nucleation rates seem independent of
649 HOM concentrations, implying that these mechanisms may occur in competition with one another.
650 The HOMs present in this study occur mostly from oxidation of alkylbenzenes and
651 monoterpenes, SVOC arose from mostly isoprene and alkylbenzene oxidation, whereas LVOC and
652 ELVOC arose from alkylbenzene, monoterpene and PAH oxidation together, with a strong
653 dependence of their concentration on both temperature, and VOC concentration. Concentrations of
654 species associated with coastal and oceanic sources such as MSA and HIO₃ were low, even though
655 NPF was consistently associated with southerly and south-westerly air masses, indicating that
656 anthropogenic emissions of SO₂, (largely arising from shipping emissions in Barcelona) and
657 aromatic organic compounds are more important precursors for compounds which initiate new
658 particle formation and growth than oceanic emissions. High HOM signals are seen to be a necessary
659 condition for new particle growth past 10 nm, with the most significant difference between days
660 with and without particle growth being SVOC concentrations (factor of 2 difference), while
661 modelled growth rates from condensation of these organic compounds, alongside H₂SO₄, MSA and

662 HIO₃ were shown to match growth rates within measurement error and it is evident that HOM
663 concentrations are dependent upon temperature and VOC precursor concentration. Specifically
664 large HOMs of extremely low volatility >500 m/Q are absent in the mass spectra on these days
665 without particle formation, as well as particle formation without growth Thus, oxidation of traffic
666 derived alkylbenzenes and PAHs, and to a lesser degree, isoprene and monoterpene emissions is a
667 significant determinant of new particle growth in this environment.-

668

669 These results are consistent with extensive chamber and flow tube studies on particle formation
670 from sulphuric acid, amines and HOMs, and further, nucleation rates relative to sulphuric acid are
671 similar to many tropospheric observations. Barcelona is representative of many Mediterranean
672 urban environments, with moderate pollution, influence of shipping emissions, and high insolation,
673 and ~~thus~~ the present study reveals the complexity of NPF mechanisms in these environments.

674

675 **DATA AVAILABILITY**

676 Data supporting this publication are openly available from the UBIRA eData repository at
677 <https://doi.org/10.25500/edata.bham.00000434>

678

679 **AUTHOR CONTRIBUTIONS**

680 RMH and XQ conceived the study, JB and DCSB carried out the [CI-APi-TOF](#) and related
681 measurements with assistance from AA and MCM. The VOC measurements were proposed by NM
682 and collected by BT-R. JB wrote the first draft of the manuscript which was enhanced by
683 contributions from the co-authors.

684

685 **COMPETING INTERESTS**

686 The authors have no conflict of interests.

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694

695 **FIGURE LEGENDS:**

696

697 **Figure 1:** Average SMPS contour plots for (a) non-event days, (b) full-event days and (c) burst-
698 event days.

699

700 **Figure 2:** Box plots for days of non-event, full-event and burst-event, showing (a) condensation
701 sink, (b) temperature, and (c) global radiation from hourly data. “Full-event” and “burst-
702 event” include data across the entire day.

703

704

705 **Figure 3:** Box plots for days of non-event, full-event and burst-event, showing (a) sulphuric acid,
706 (b) C₂ and C₄ amines, as clustered with the nitrate dimer and trimer, and (c) summed
707 HOM concentration from C₅₊ from hourly data. Units for ammonia + amines are
708 normalised counts, as no calibration was performed. Event days include data across the
709 full event day.

710

711 **Figure 4:** Formation rate (J_{1.9}) plotted against sulphuric acid monomer concentration, coloured by
712 condensation sink. Circles represent burst-events, squares represent full events. Data is
713 for hourly averages across NPF periods, typically within the hours 08:00 – 16:00. Slope
714 of the line = $4.9 \cdot 10^{-5} \text{ s}^{-1}$. Error bars represent systematic uncertainties on [H₂SO₄] and
715 J_{1.9}.

716

717 **Figure 5:** Formation rate plotted against sulphuric acid monomer concentration for data collected
718 from Barcelona. Tan circles represent burst-events, purple squares represent full events.
719 as well as that for the H₂SO₄-H₂O (blue inverted triangles), H₂SO₄-NH₃-H₂O (yellow
720 inverted triangles), H₂SO₄-DMA-H₂O (pink triangles), and H₂SO₄-BioOxOrg-H₂O
721 (brown diamonds) systems from the CLOUD chamber (Kürten et al., 2018 Kirkby et al.,
722 2011; Riccobono et al., 2014). CLOUD chamber experiments were performed at 278 K
723 and 38 – 39 % RH. Data is for hourly averages across NPF periods, typically within the
724 hours 08:00 – 16:00. Error bars represent systematic uncertainties on [H₂SO₄] and J_{1.9}.

725

726 **Figure 6:** Sulphuric acid dimer concentration plotted against monomer concentration, showing
727 burst-event periods (tan circles), full event periods (purple squares), non-event periods
728 (green inverted triangles), and the ratio of sulphuric acid dimer:monomer in the CLOUD
729 chamber for the H₂SO₄-H₂O-DMA system (pink triangles) (Almeida et al., 2013).
730 Dashed line represents the dimer concentration produced by ion induced clustering in the
731 chemical ionization unit (Zhao et al., 2010). CLOUD chamber experiments were
732 performed at 278 K and 38 – 39 % RH. Data is for hourly averages across NPF periods,
733 typically within the hours 08:00 – 16:00. Error bars represent systematic uncertainties on
734 [H₂SO₄] and [(H₂SO₄)₂].

735

736

737 **Figure 7:** Influencing factors on HOM concentration, showing (a) C₅₋₁₀ HOM concentration plotted
738 against temperature, coloured by global radiation. Ellipsis shows 95% confidence on a
739 multivariate t-distribution. (b) HOM concentration by carbon number plotted against
740 parent VOC mixing ratio. These are segregated by carbon number/VOC, i.e. C₇ HOMs
741 plotted against toluene, under the assumption that toluene oxidation is the main producer
742 of C₇ HOMs. Time for both plots is of hourly time resolution.

743

744 **Figure 8:** Concentrations of all oxygenated organic molecules and HOMs binned to integer
745 $\text{Log}_{10}(\text{C}^*)$ values, coloured by DBE.

746
747 **Figure 9:** Mass defect plots for (a) non-event, (b) full-event, and (c) burst-event periods, data
748 taken from 10:00 – 15:00 on the days 11/07/2018, 16/07/2018 and 15/07/2018
749 respectively. Size corresponds to mass spectral peak area. Ions are coloured according
750 to identified chemical composition. *Blue* points correspond to HOMs containing all
751 organic species with ≥ 5 carbon atoms and ≥ 6 oxygen atoms, and an O:C ratio of >0.6 .
752 *Purple* points correspond to the same but for species containing 1-2 nitrogen atoms.
753 Species not meeting this HOM criteria were classed generally as OVOCs which are
754 coloured *brown*, with the nitrogen containing OVOCs coloured *orange*. Sulphur acids
755 (*red*) include ions HSO_4^- , CH_3SO_3^- and SO_5^- , as well as the sulphuric acid dimer. Iodine
756 acids (*green*) contains both IO_3^- and I^- (the latter presumably deprotonated hydrogen
757 iodide). Unidentified points are left uncoloured.

759

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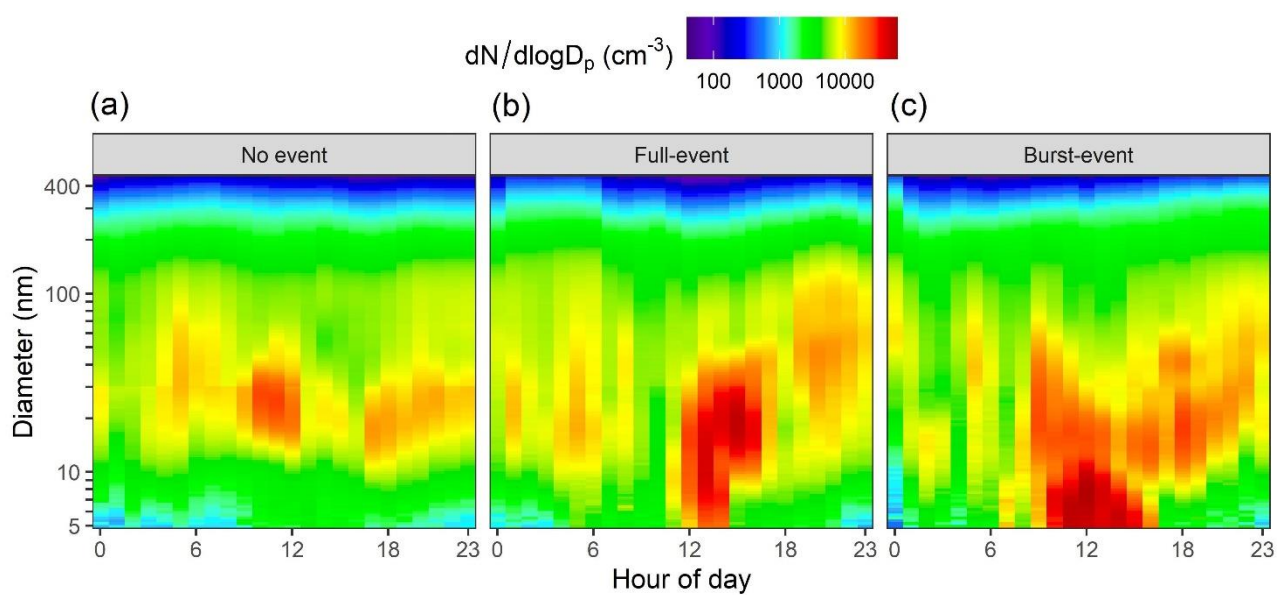
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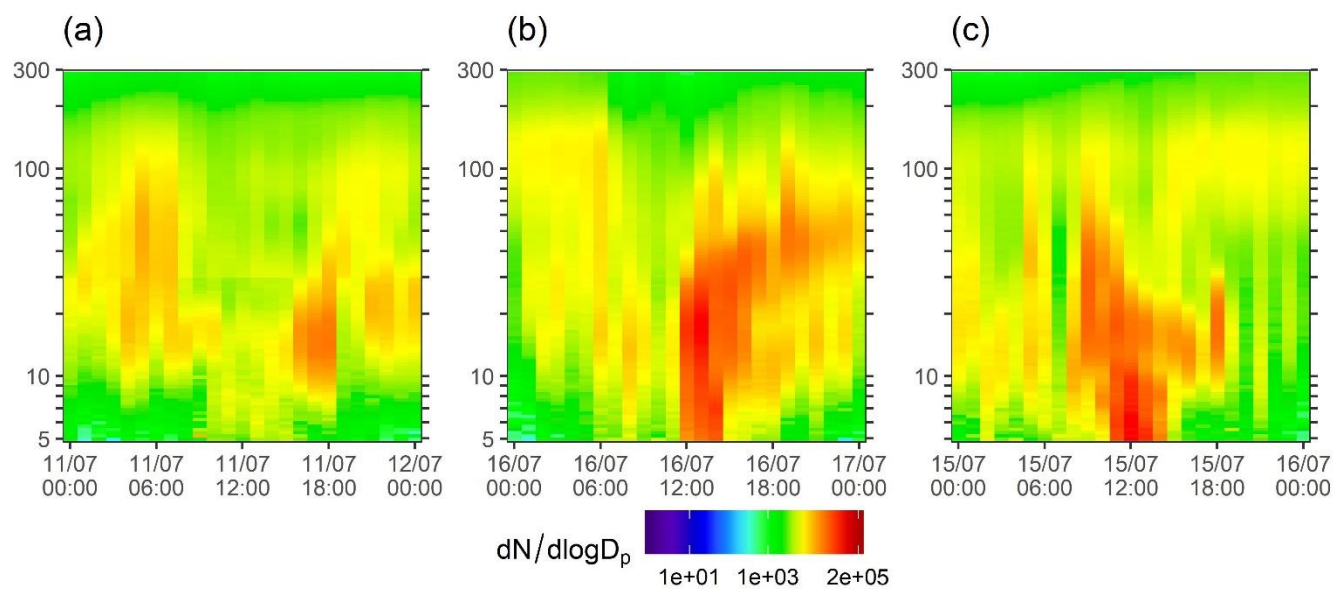
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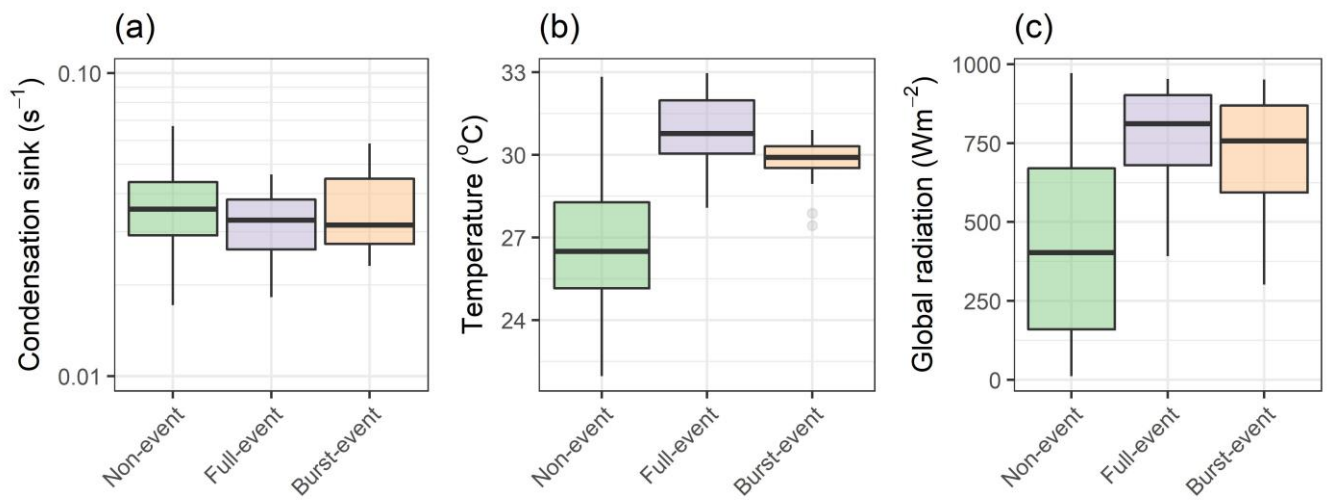


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1279 **Figure 1:** Example Average SMPS contour plots of for (a) non-event days, (b) full-event days and
 1280 (c) burst-event days.

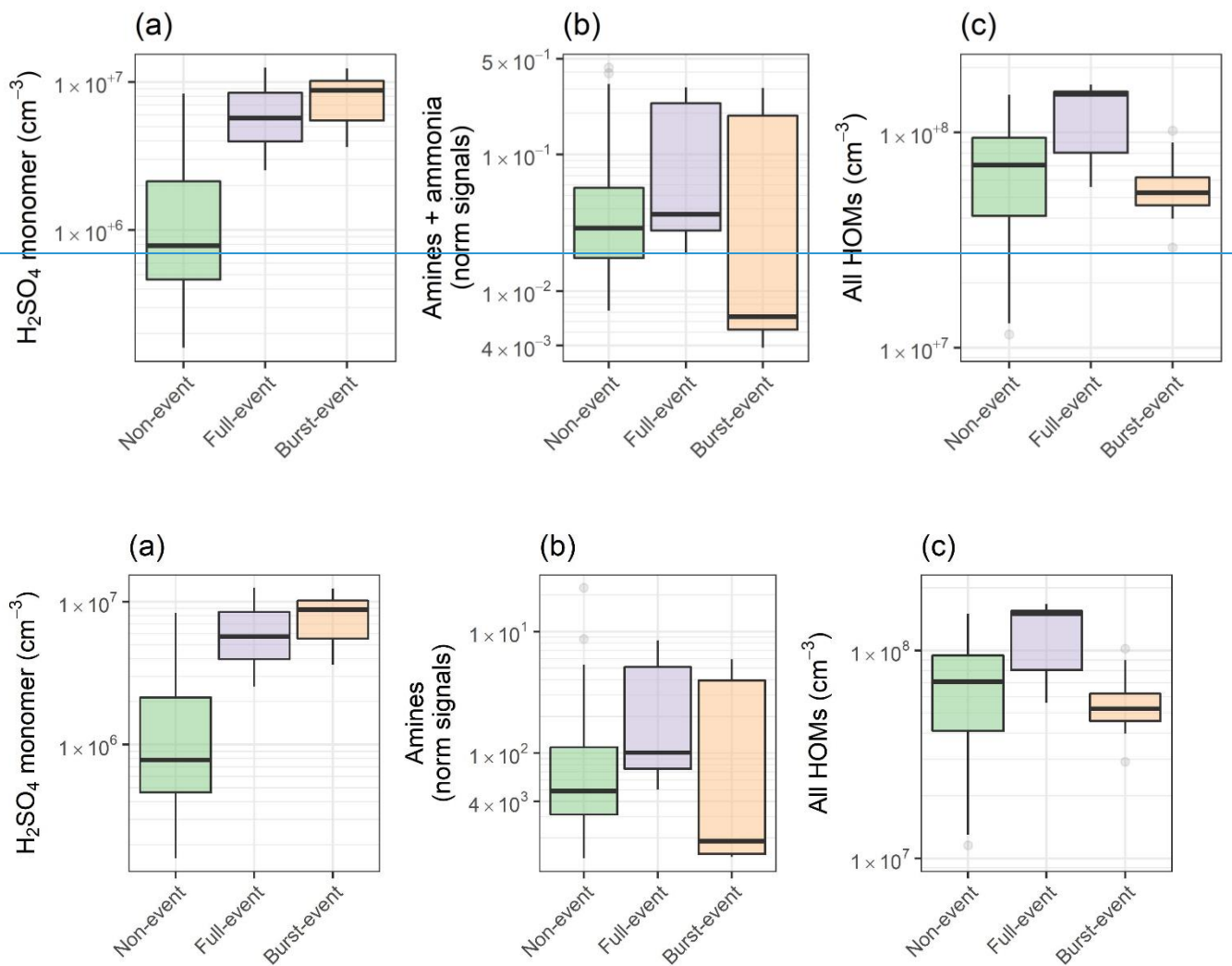
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1283 **Figure 2:** Box plots for days of non-event, full-event and burst-event, showing (a) condensation sink,
 1284 (b) temperature, and (c) global radiation [from hourly data](#). [“Full-event” and “burst-event” include](#)
 1285 [data across the entire day](#).

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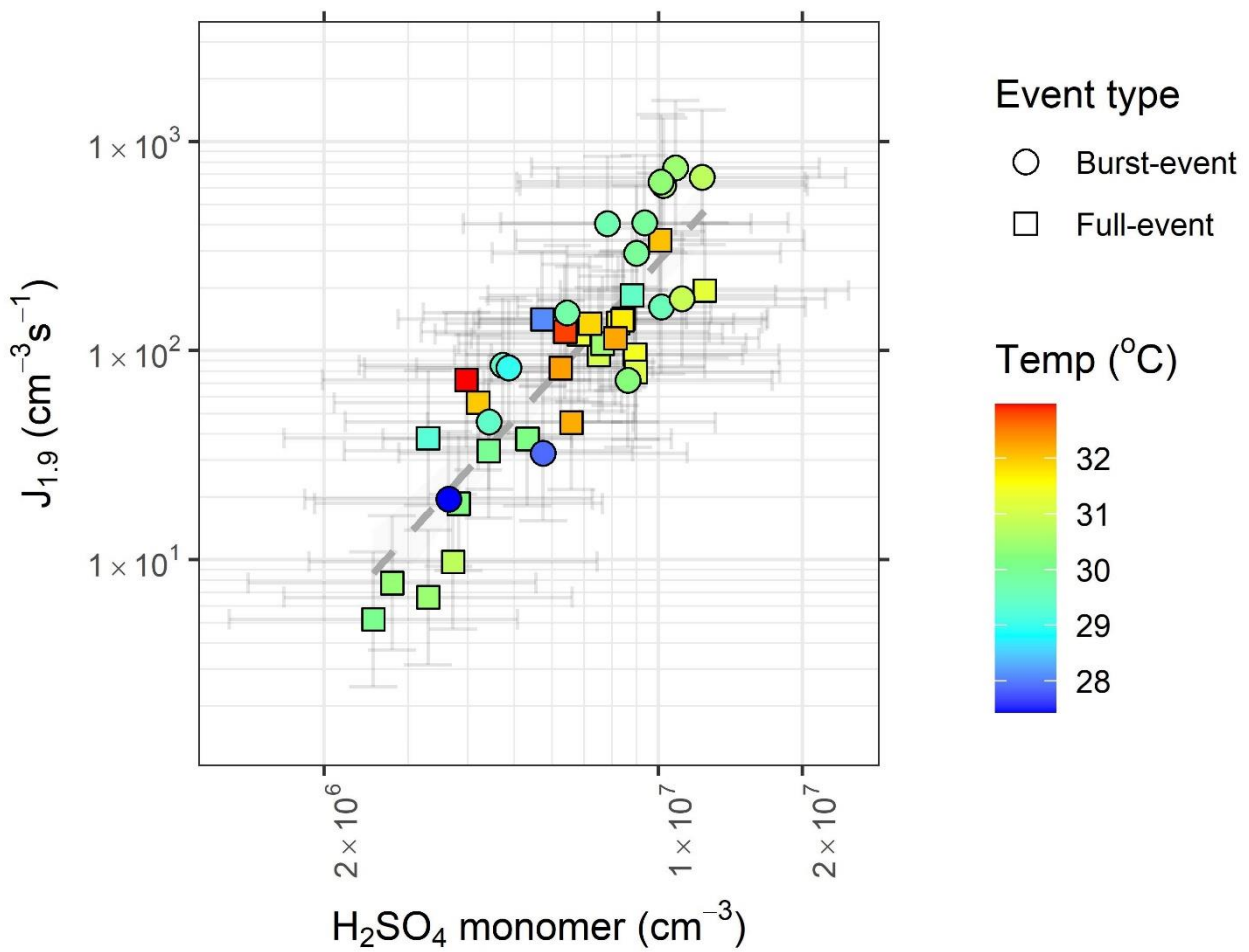
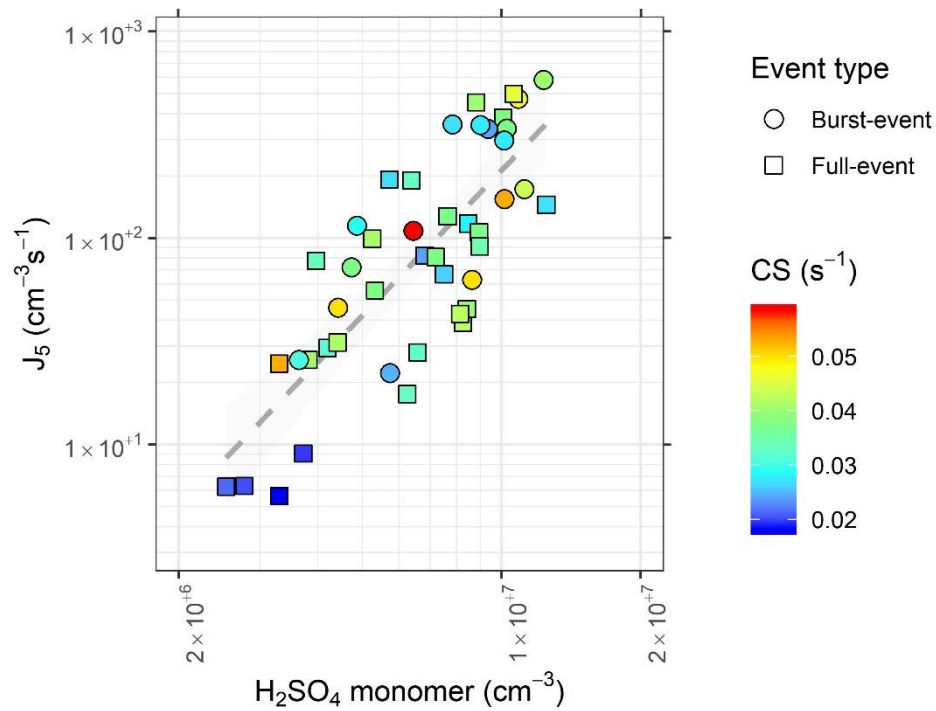


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1289 **Figure 3:** Box plots for days of non-event, full-event and burst-event, showing (a) sulphuric acid, (b)
 1290 [ammonia](#), C₂ and C₄ amines, as clustered with the nitrate dimer and trimer, and (c) summed HOM
 1291 concentration from C₅+ [from hourly data](#). Units for ammonia + amines are normalised counts, as no
 1292 calibration was performed. [Event days include data across the full event day.](#)

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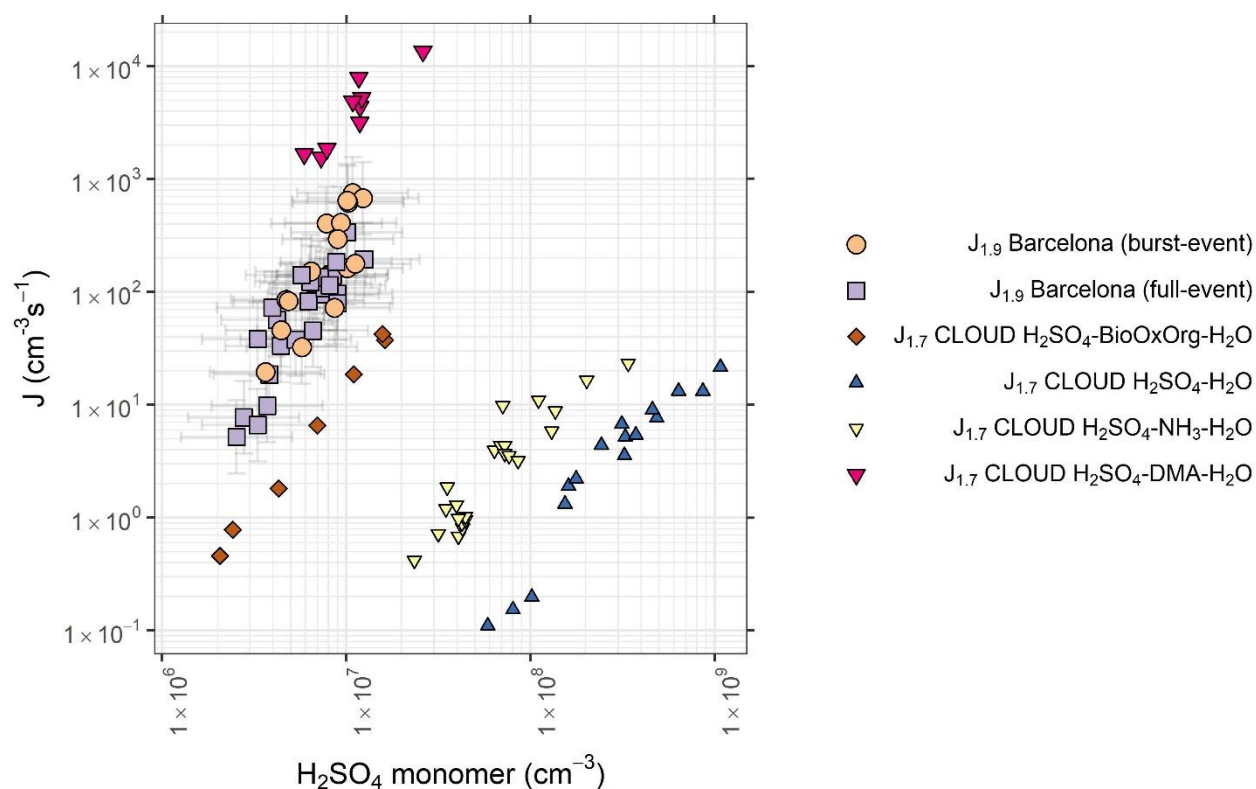
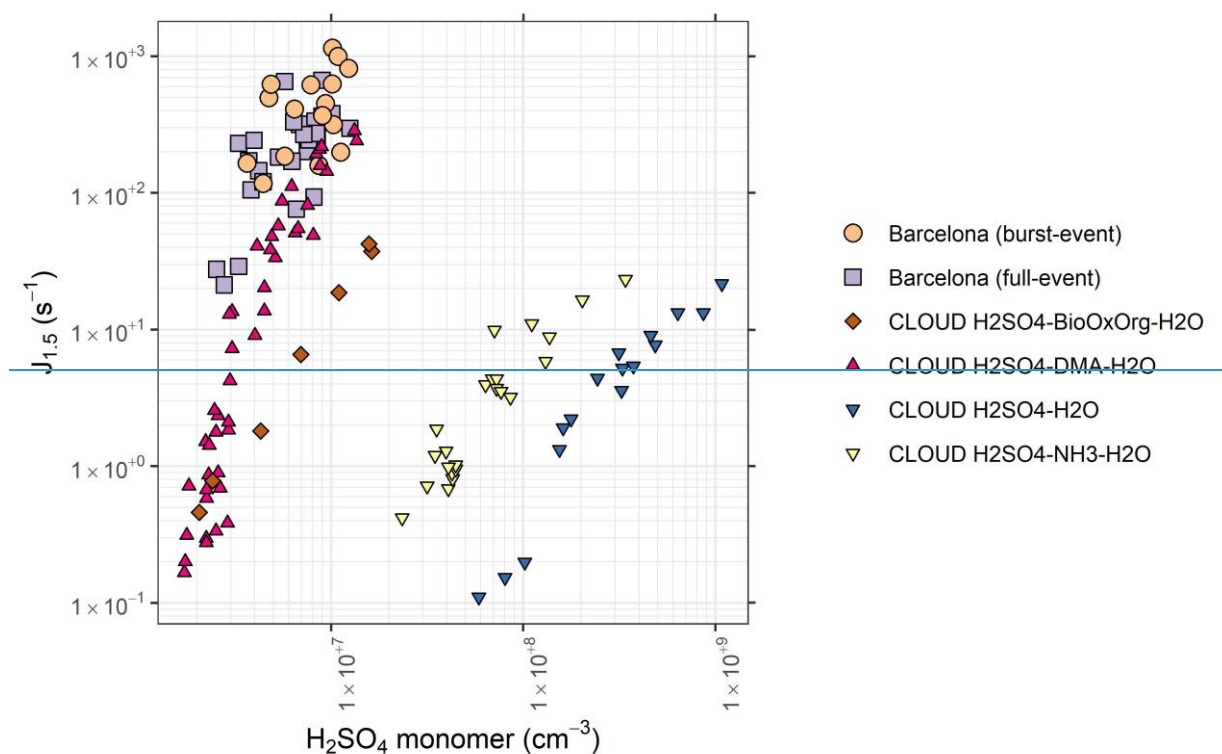


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1296 **Figure 4:** Formation rate ($J_5 J_{1.9}$) plotted against sulphuric acid monomer concentration, coloured by
 1297 condensation sink. Circles represent burst-events, squares represent full events. [Data is for hourly](#)

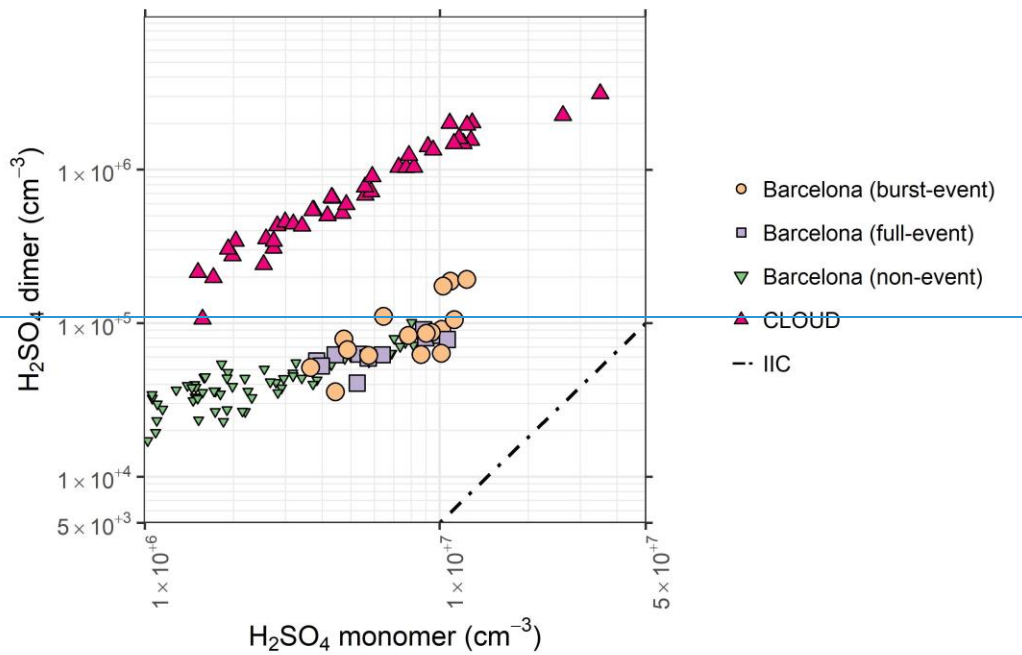
1298 averages across NPF periods, typically within the hours 08:00 – 16:00. Slope of the line = $4.9 \cdot 10^{-5} \text{ s}^{-1}$
1299 ¹. Error bars represent systematic uncertainties on $[\text{H}_2\text{SO}_4]$ and $J_{1.9}$
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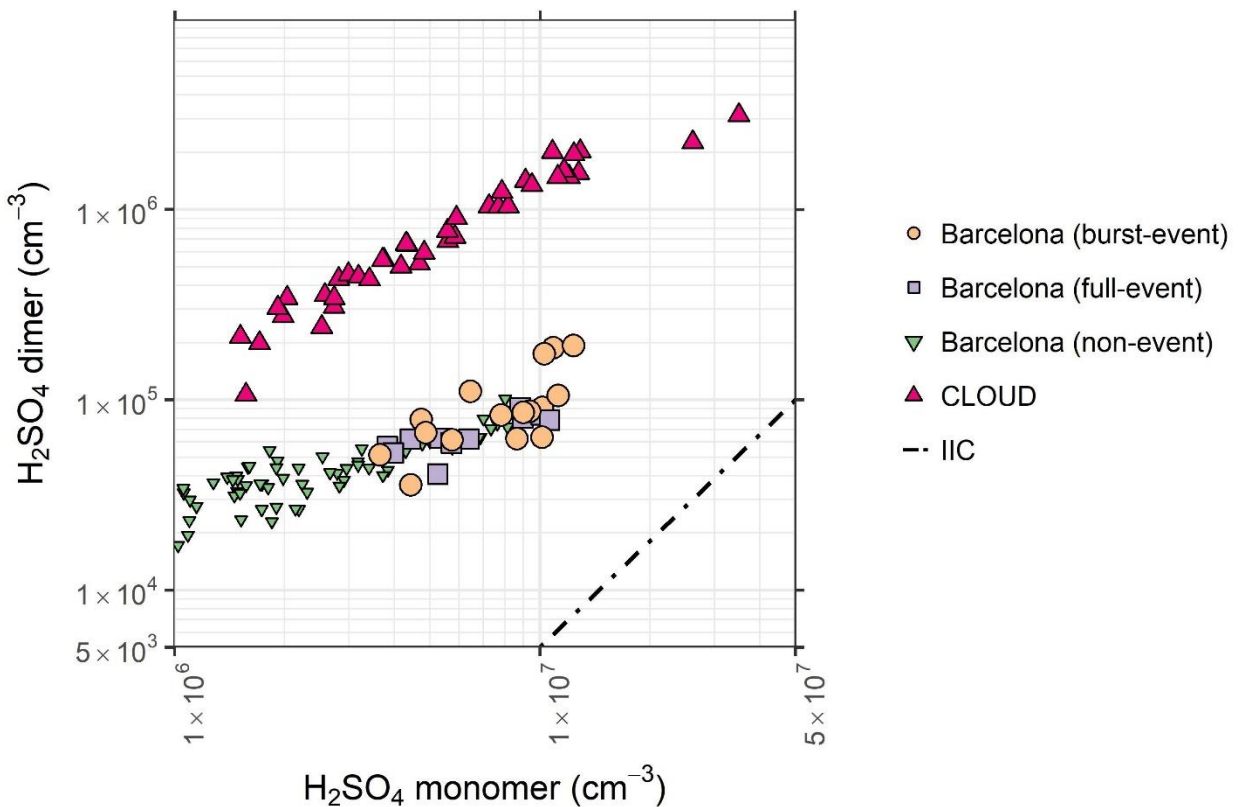
1304 Figure 5: **Formation rate ($J_{1.5}$) plotted against sulphuric acid monomer concentration for data**
 1305 **collected** from Barcelona. Tan circles represent burst-events, purple squares represent full events.
 1306 as well as that for the $\text{H}_2\text{SO}_4\text{-H}_2\text{O}$ (blue inverted triangles), $\text{H}_2\text{SO}_4\text{-NH}_3\text{-H}_2\text{O}$ (yellow inverted
 1307 triangles), $\text{H}_2\text{SO}_4\text{-DMA-H}_2\text{O}$ (pink triangles), and $\text{H}_2\text{SO}_4\text{-BioOxOrg-H}_2\text{O}$ (brown diamonds)
 1308 systems from the CLOUD chamber ([Almeida et al., 2013](#); [Kürten et al., 2018](#) Kirkby et al., 2011;

1309 Riccobono et al., 2014). CLOUD chamber experiments were performed at 278 K and 38 – 39 %
1310 RH. Data is for hourly averages across NPF periods, typically within the hours 08:00 – 16:00. Error
1311 bars represent systematic uncertainties on [H₂SO₄] and J_{1,9}.
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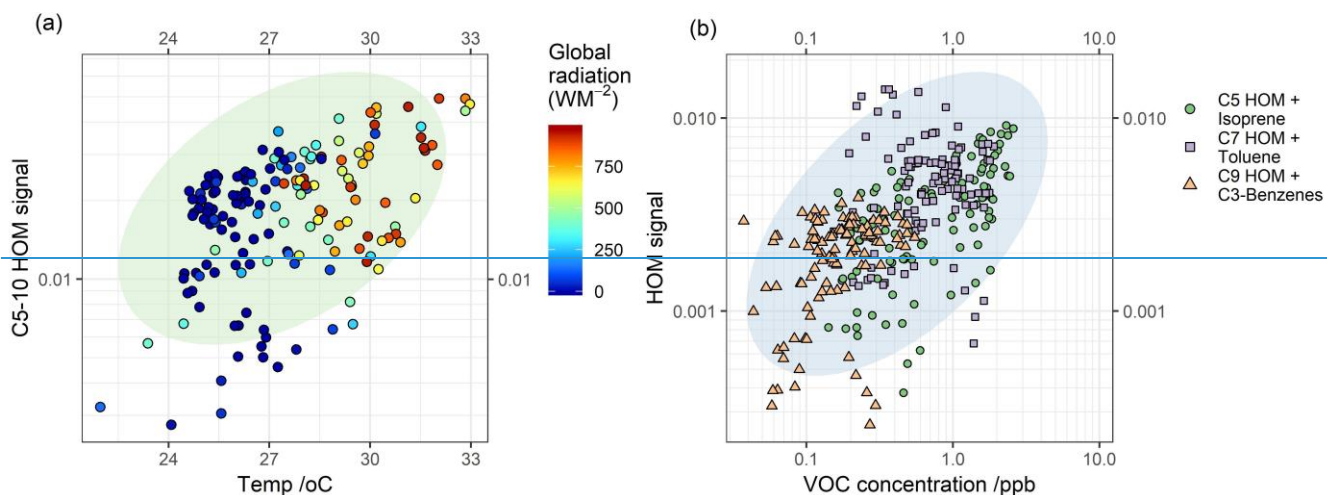
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1316 **Figure 6:** Sulphuric acid dimer concentration plotted against monomer concentration, showing burst-
 1317 event periods (tan circles), full event periods (purple squares), non-event periods (green inverted
 1318 triangles), and the ratio of sulphuric acid dimer:monomer in the CLOUD chamber for the H_2SO_4 -
 1319 H_2O -DMA system (pink triangles) (Almeida et al., 2013). Dashed line represents the [lower limit of](#)
 1320 dimer concentration produced by ion induced clustering in the chemical ionization unit (Zhao et al.,
 1321 2010). [CLOUD chamber experiments were performed at 278 K and 38 – 39 % RH. Data is for hourly](#)

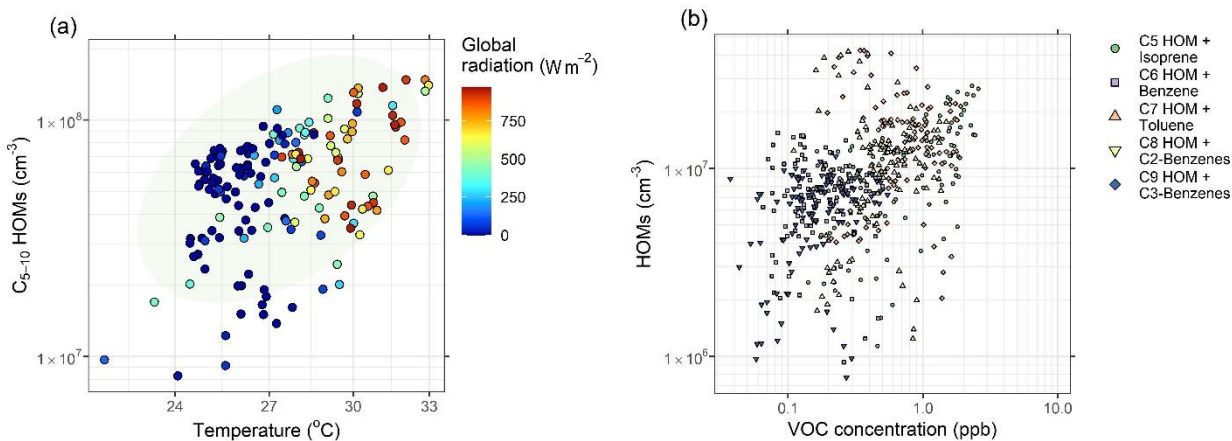
1322 averages across NPF periods, typically within the hours 08:00 – 16:00. Error bars represent systematic
1323 uncertainties on $[\text{H}_2\text{SO}_4]$ and $[(\text{H}_2\text{SO}_4)_2]$.

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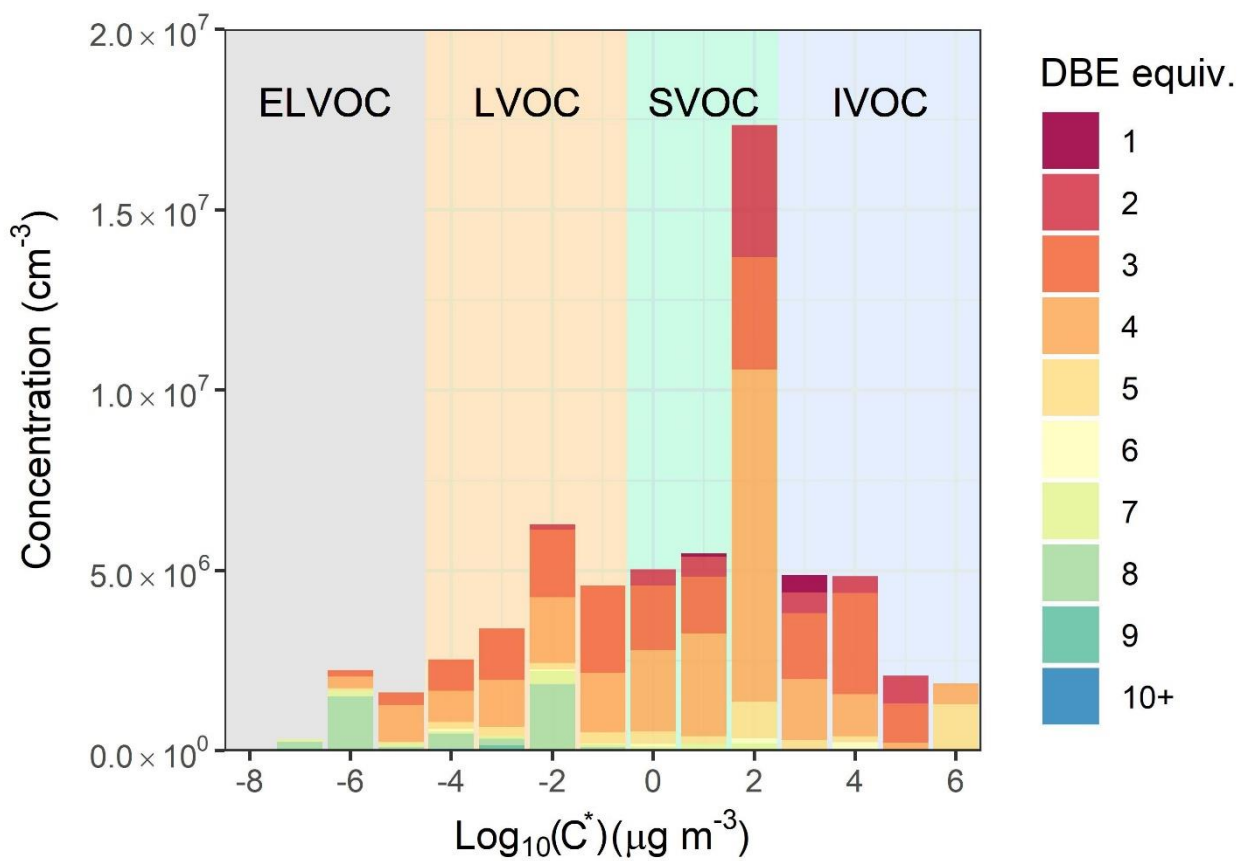


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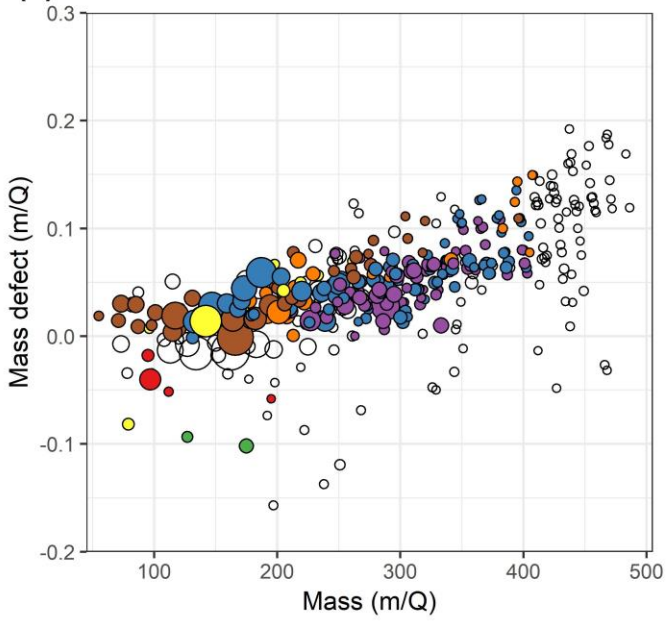
1328 **Figure 7:** Influencing factors on VOC-HOM concentration, showing (a) temperature plotted against
 1329 C₅₋₁₀ HOM signal concentration plotted against temperature, coloured by global radiation. Ellipses
 1330 shows 95% confidence on a multivariate t-distribution. and (b) HOM concentration by carbon number
 1331 plotted against parent VOC mixing ratio. VOC concentration plotted against HOM signal. These are
 1332 segregated by carbon number/VOC, i.e. C₇ HOMs plotted against toluene, under the assumption that
 1333 toluene oxidation is the main producer of C₇ HOMs. Time for both plots is of hourly time
 1334 resolution. ~~Ellipses show 95% confidence on a multivariate t-distribution.~~



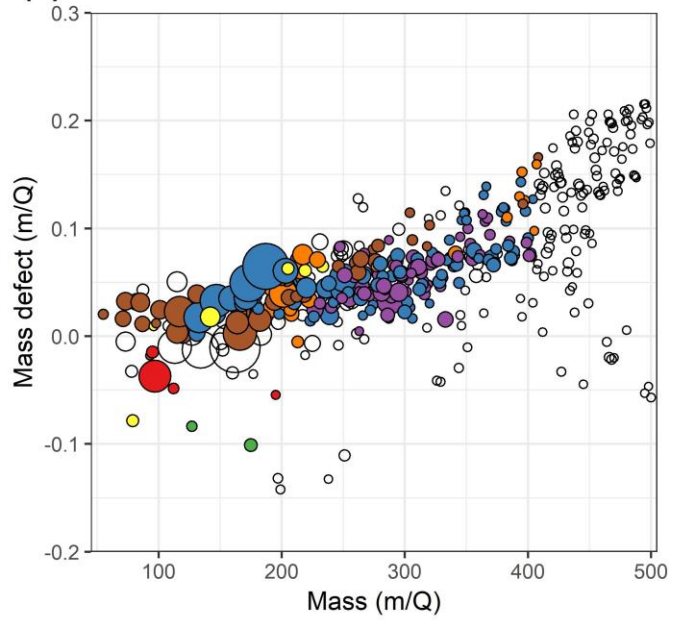
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1336 **Figure 8:** Concentrations of all oxygenated organic molecules and HOMs binned to integer $\text{Log}_{10}(C^*)$
 1337 values, coloured by DBE.

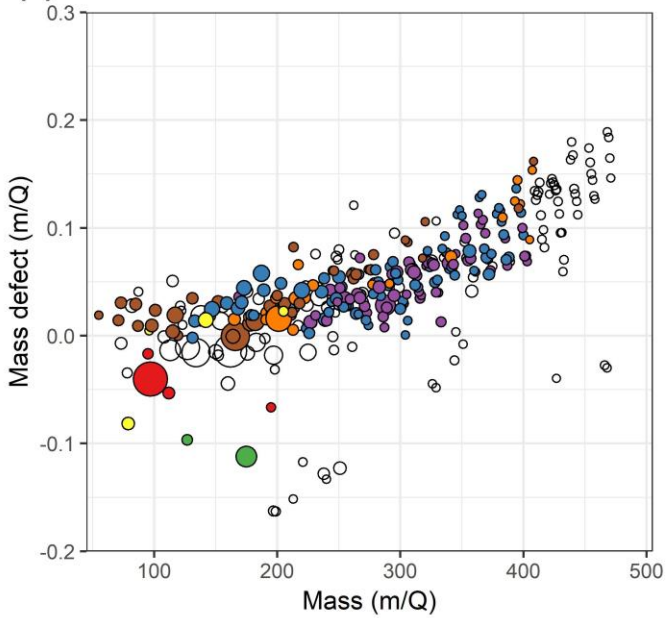
(a) Non-event



(b) Full-event



(c) Burst-event

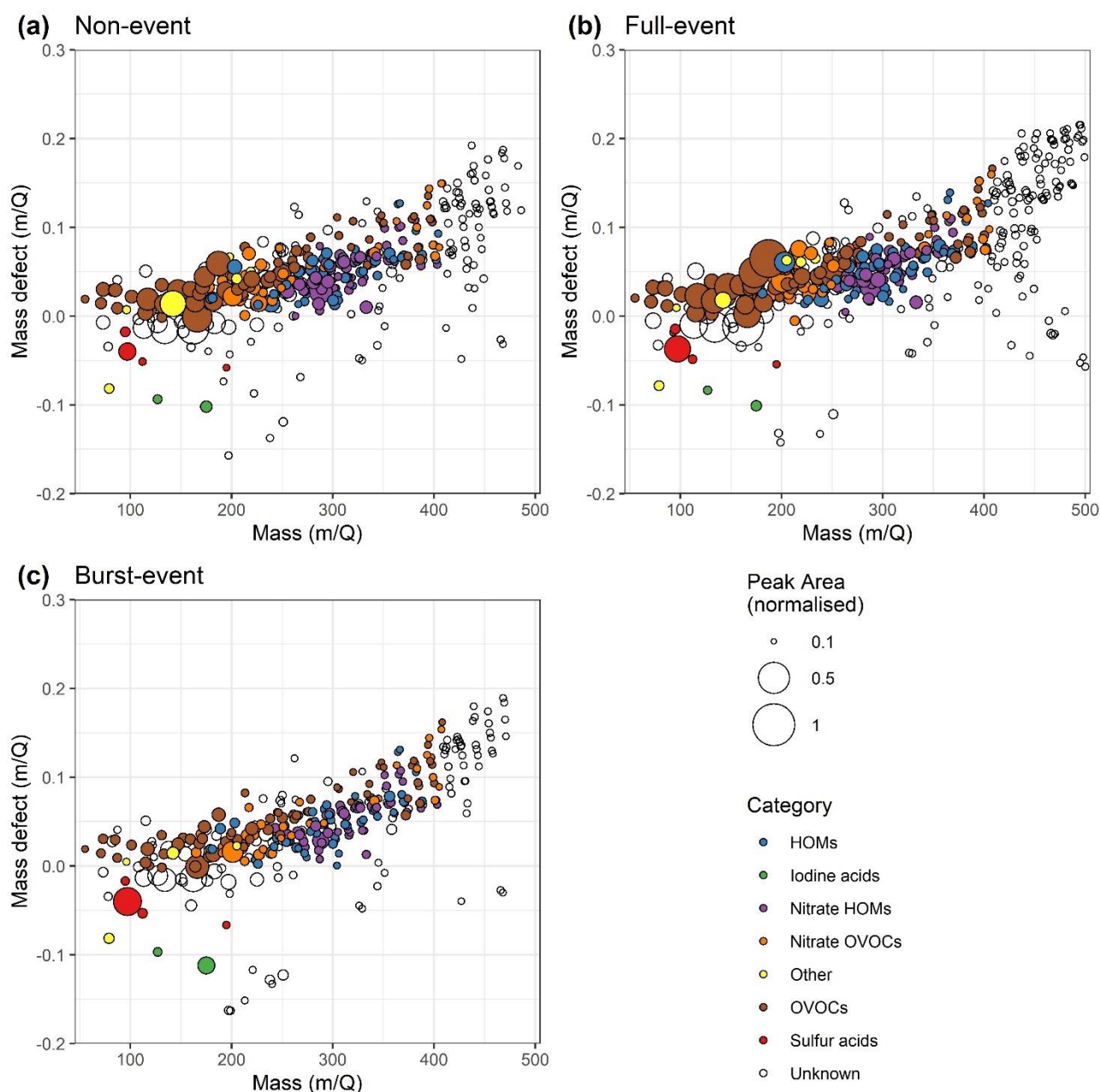


Category

- HOMs
- Iodine acids
- Nitrate HOMs
- Nitrate OVOCs
- Other
- OVOCs
- Sulfur acids
- Unknown

Peak Area (normalised)

- 0.1
- 0.5
- 1



1339

1340 **Figure 89:** Mass defect plots for (a) non-event, (b) full-event, and (c) burst-event periods, data taken
 1341 from 10:00 – 15:00 on the days 11/07/2018, 16/07/2018 and 15/07/2018 respectively. Size
 1342 corresponds to mass spectral peak area. Ions are coloured according to identified chemical
 1343 composition. *Blue* points correspond to HOMs containing all organic species with ≥ 5 carbon atoms
 1344 and ≥ 6 oxygen atoms, and an O:C ratio of >0.6 . *Purple* points correspond to the same but for species
 1345 containing 1-2 nitrogen atoms. Species not meeting this HOM criteria were classed generally as
 1346 OVOCs which are coloured *brown*, with the nitrogen containing OVOCs coloured *orange*. Sulphur
 1347 acids (*red*) include ions HSO_4^- , CH_3SO_3^- and SO_5^- , as well as the sulphuric acid dimer. Iodine acids
 1348 (*green*) contains both IO_3^- and I^- (the latter presumably deprotonated hydrogen iodide). Unidentified
 1349 points are left uncoloured.

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1 **Supplementary Information**

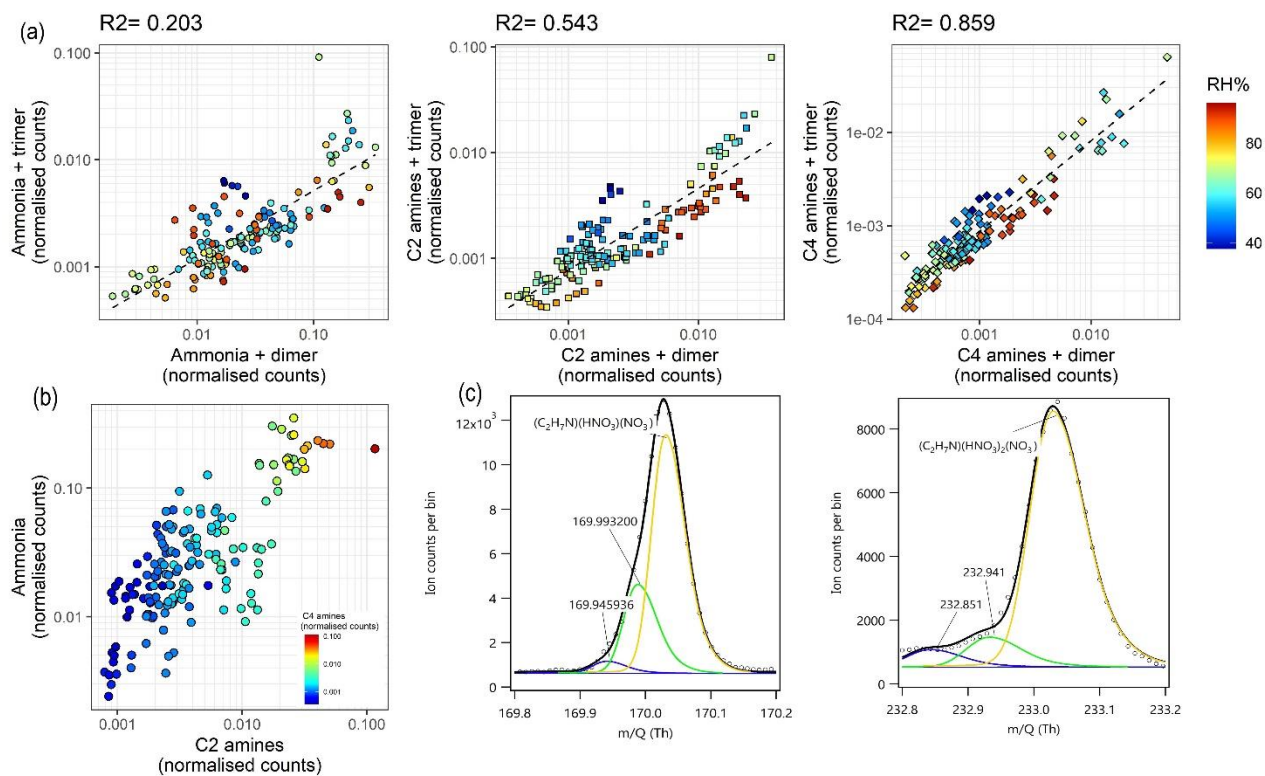
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3 **Molecular Insights into New Particle Formation in Barcelona,**
4 **Spain**

5

6 **James Brean, David C.S. Beddows, Zongbo Shi, Brice Temime-Roussel,**
7 **Nicolas Marchand, Xavier Querol, Andrés Alastuey, María Cruz**
8 **Minguillón, and Roy M. Harrison**

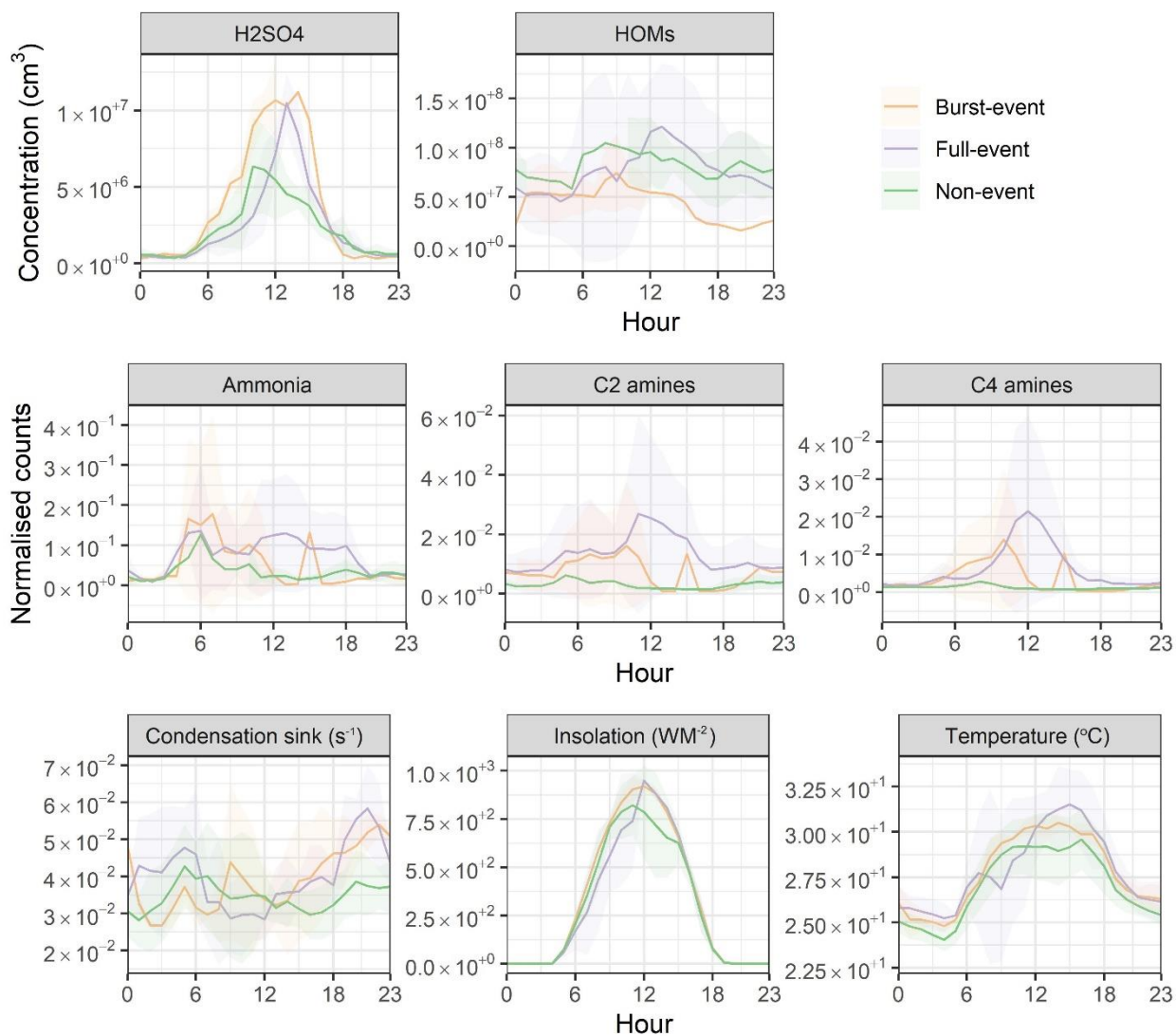
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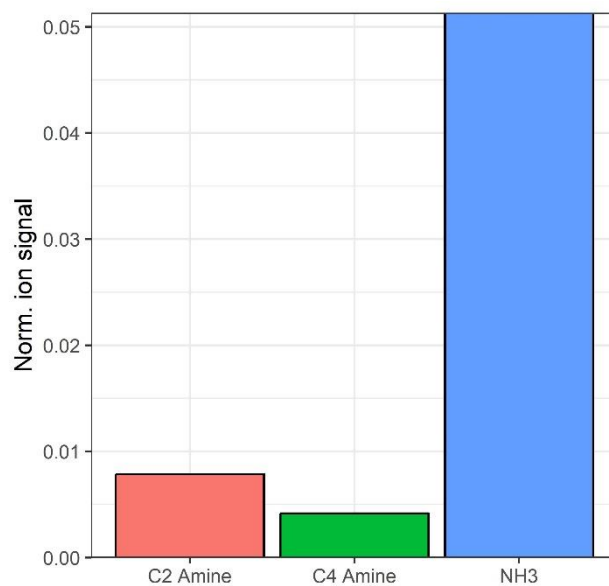
11 **Figure S1:** Ammonia and amine measurements via CI-APi-ToF, showing (a) ammonia, C₂
 12 and C₄ amines as measured clustered with the nitrate dimer and trimer. Colour scale shows an
 13 RH dependence. (b) Ammonia plotted against C₂ amines, coloured by C₄ amine concentration,
 14 and (c) peak fits for the C₂ amine ion as clustered with the nitrate dimer and trimer.

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18 **Figure S2:** Diurnal profiles of (from top left through bottom right), H₂SO₄, HOMs, NH₃, C₂
 19 amines, C₄ amines, condensation sink, insolation and temperature. Shaded regions show 1
 20 standard deviation on the mean.

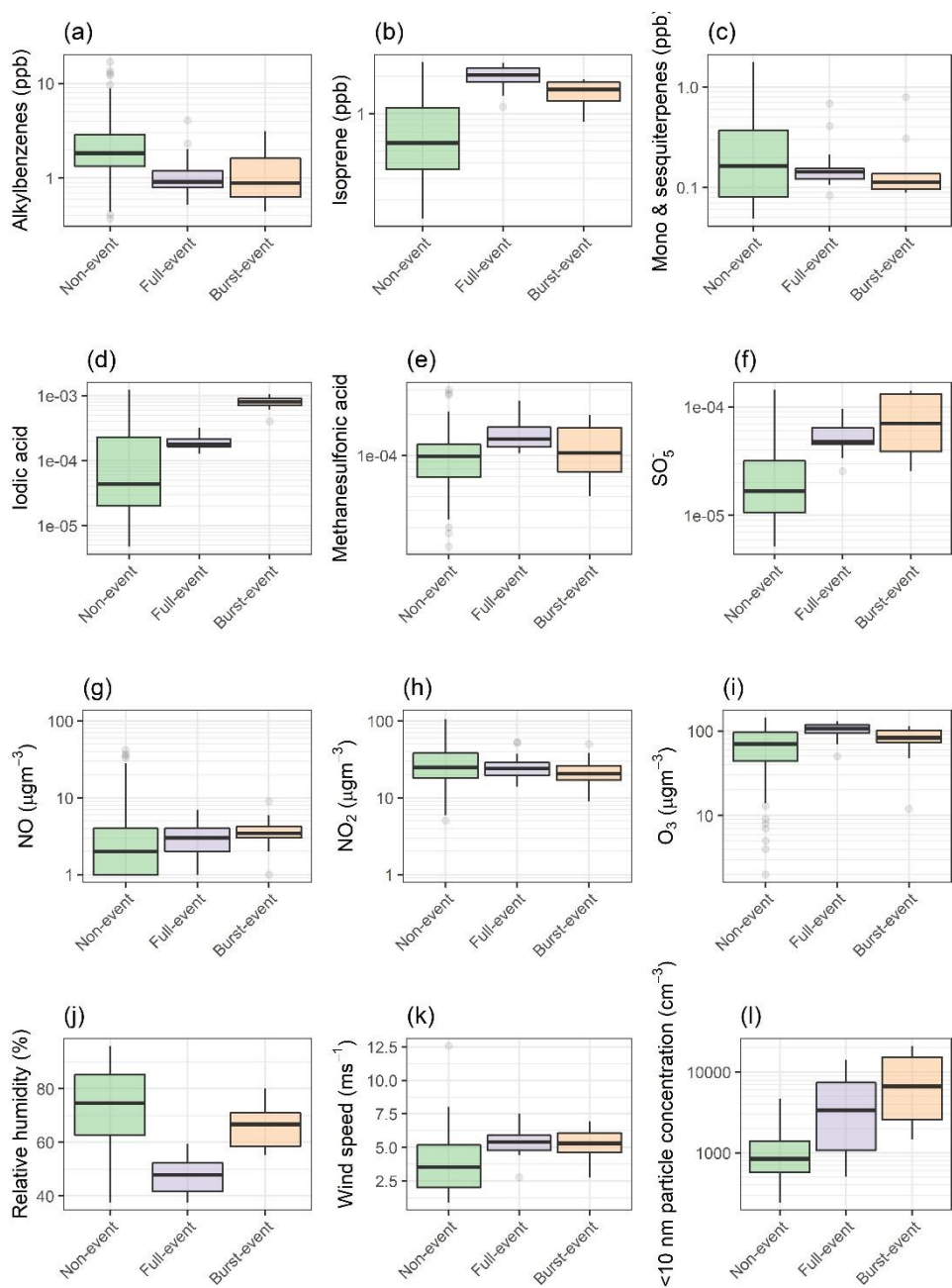


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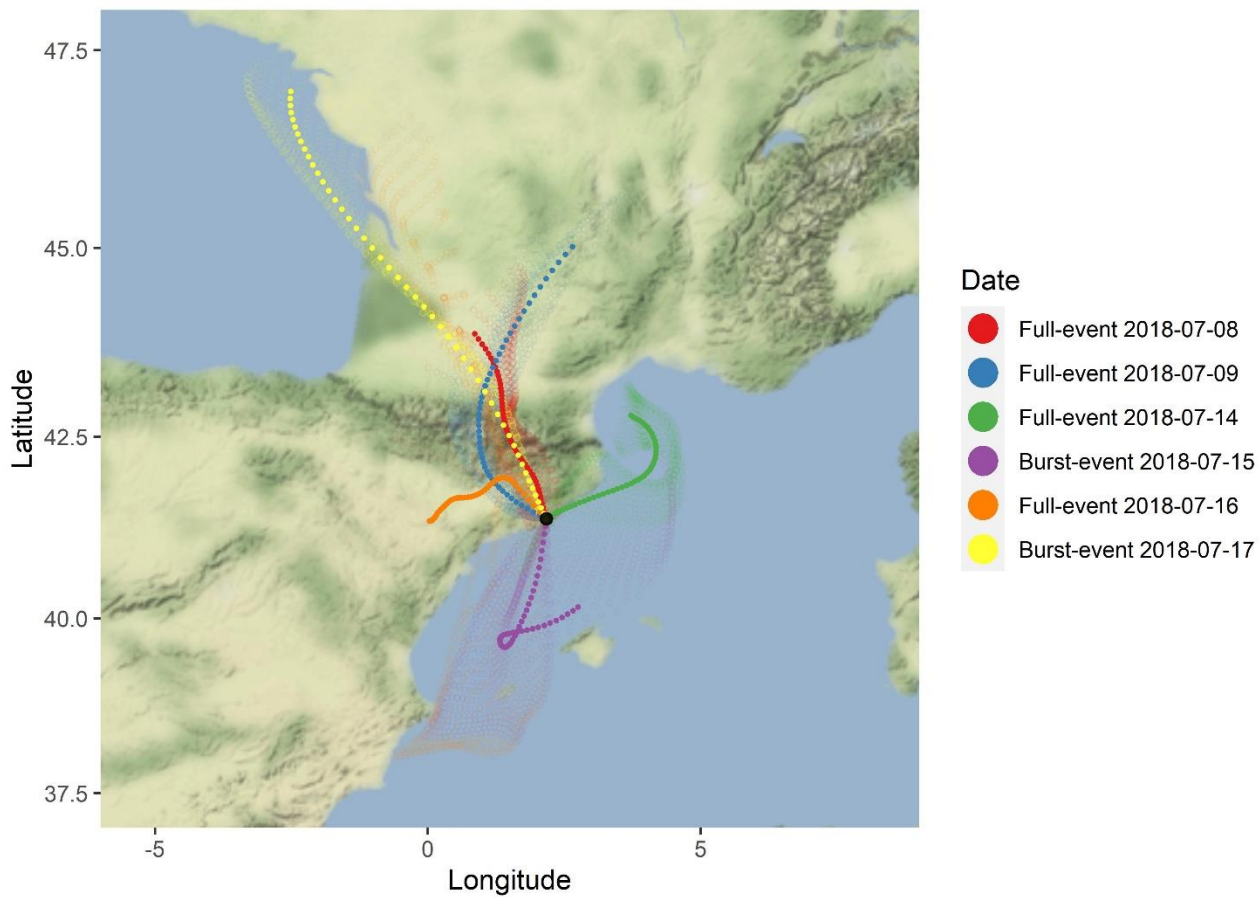
23

Figure S3: Mean ammonia and amine signals across the campaign as measured by CI-APi-ToF. Units of normalised ion counts.



24

25 **Figure S4:** Box plots as figures 2 & 3 for (a-c) VOCs as measured by PTR-ToF-MS, (d-f)
 26 other ions as measured by CI-API-ToF (units of norm. counts) (g-i) trace gases, and (j-l)
 27 meteorological and <10 nm particle count parameters.

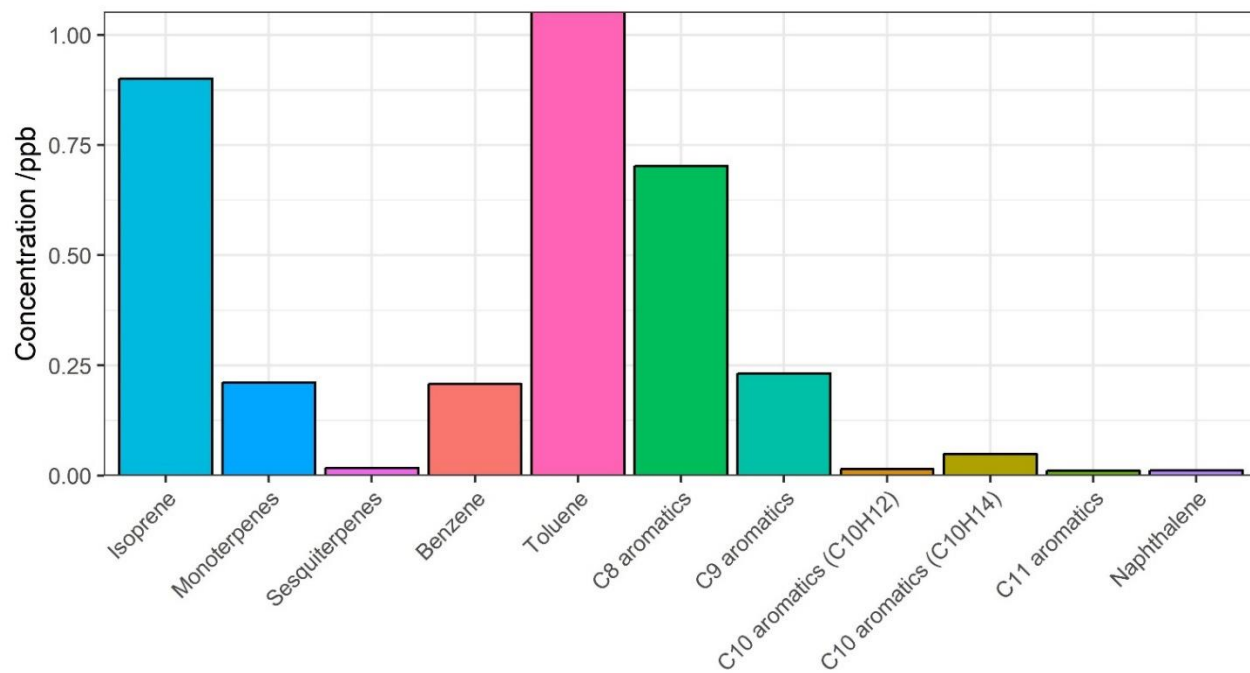


28

29 **Figure S5:** HYSPLIT 96 hour back trajectories per nucleation event. Dark dashed lines
 30 show mean trajectory per event, light dashed lines show hourly trajectories from which
 31 mean is calculated.

32

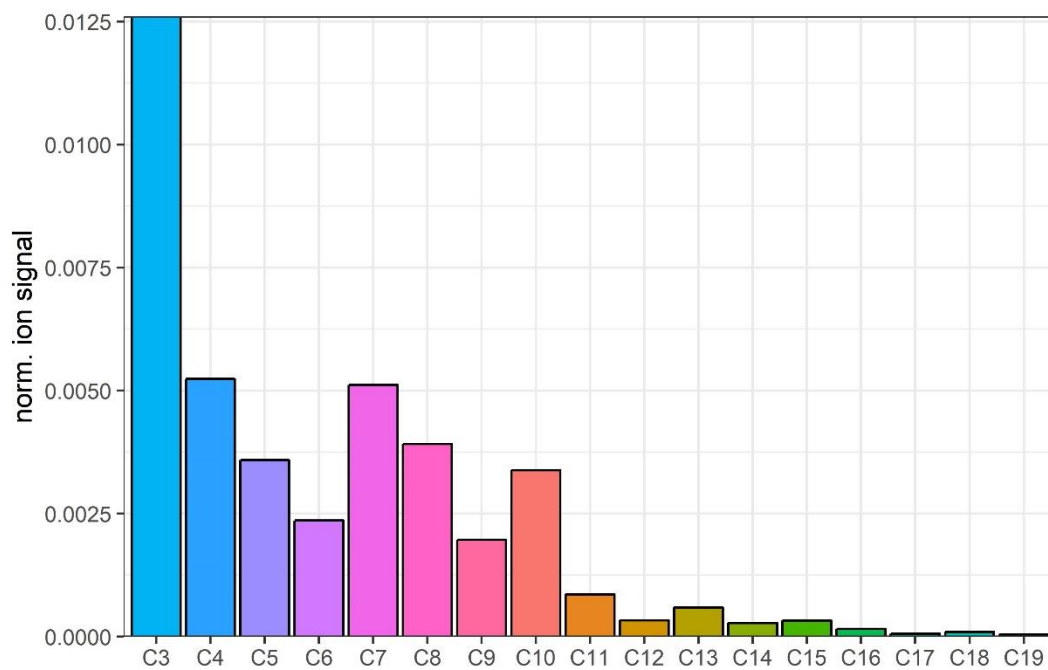
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35 **Figure S6:** Mean selected VOC concentrations across the campaign as measured by PTR-
36 ToF-MS.

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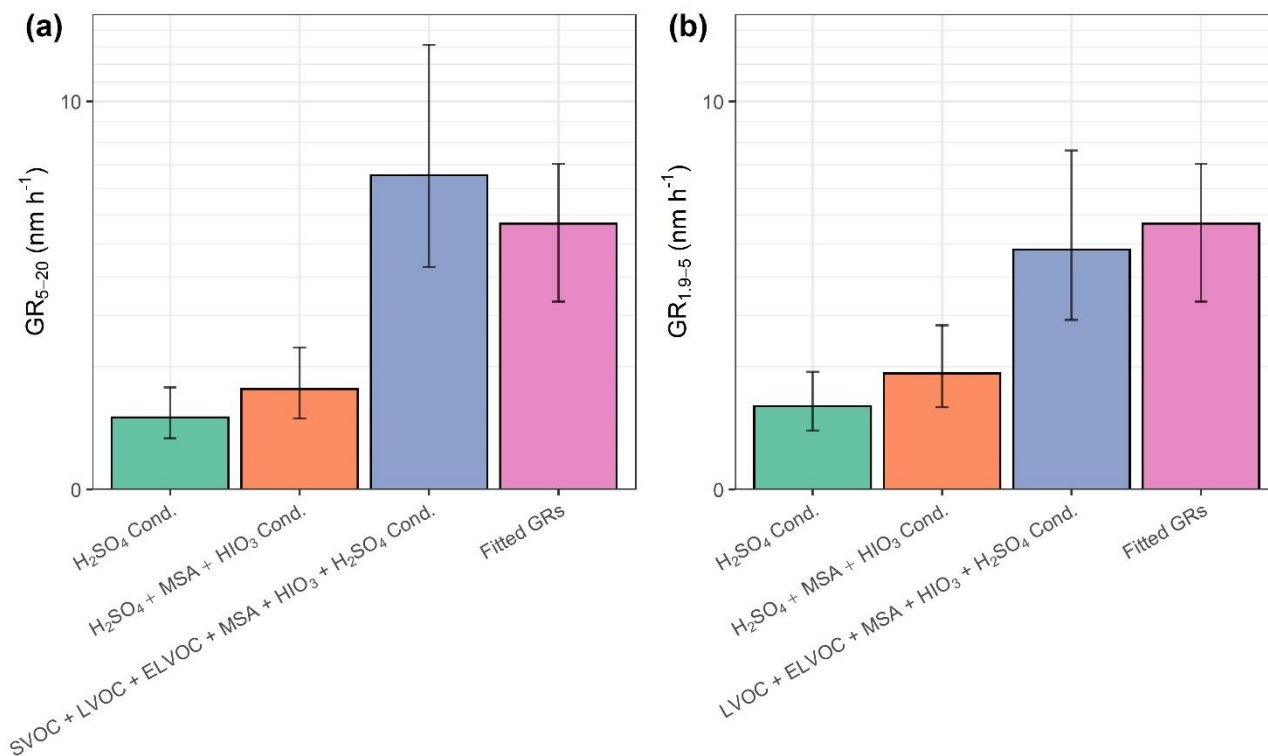
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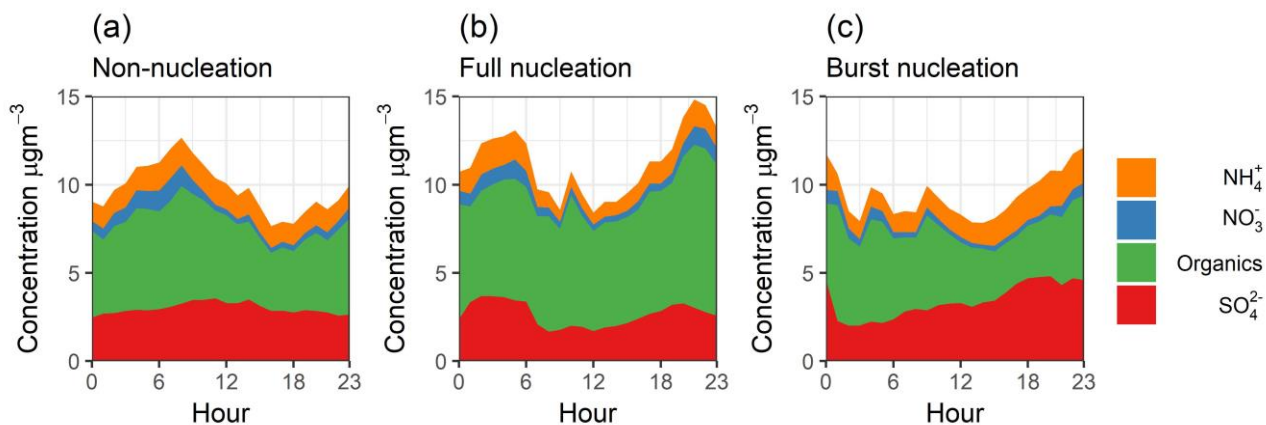
41

Figure S7: Mean ion signals per carbon number across the campaign as measured by CI-API-ToF. Units of normalised ion counts.



43

44 **Figure S8:** Condensational growth rates between (a) 5 - 20 nm and (b) 1.9 – 5 nm, calculated from
 45 H₂SO₄ condensation, H₂SO₄, MSA, and HIO₃ condensation, and SVOC, LVOC, ELVOC, H₂SO₄,
 46 MSA and HIO₃ in (a), and LVOC, ELVOC, H₂SO₄, MSA and HIO₃ in (b). Also presented are growth
 47 rates from particle count data. Error bars represent uncertainties on the concentration of species
 48 measured by CI-API-ToF, and the uncertainties from GR calculations. Systematic uncertainties from
 49 the methods of Nieminen et al. (2010) are not included.



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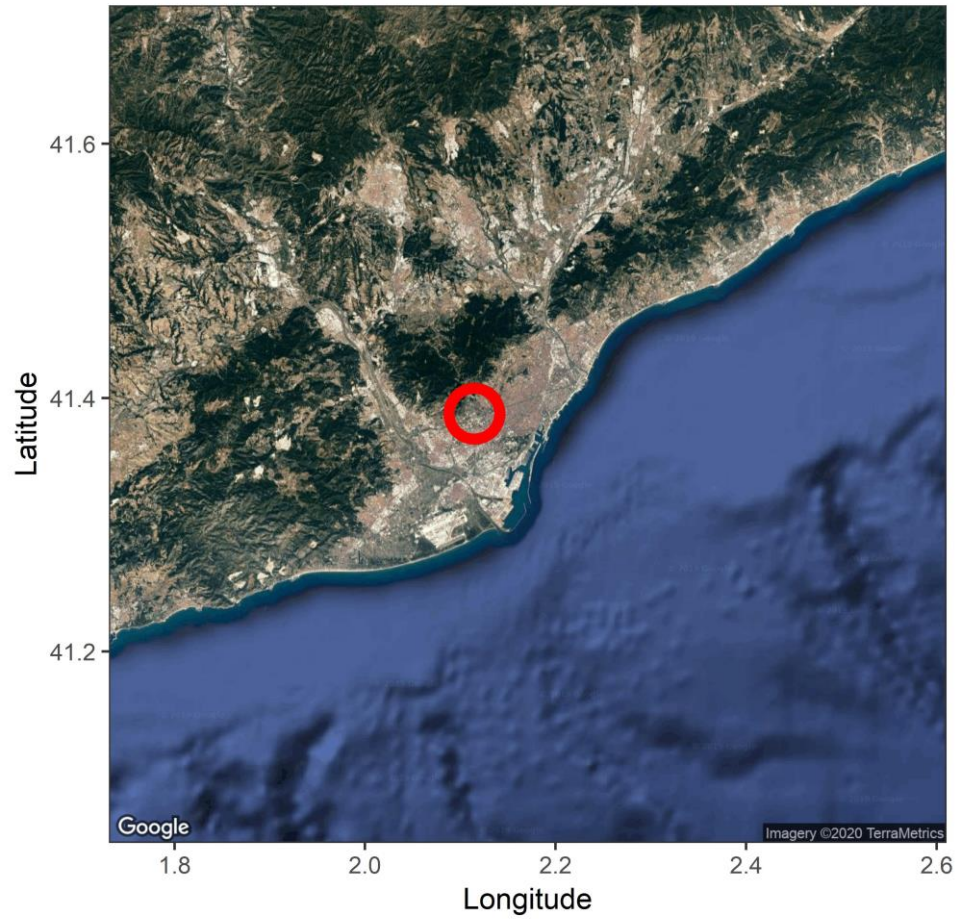
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Figure S9: Average diurnals of particle composition as measured by ACSM on (a) non-nucleation, (b) full-nucleation and (c) burst-nucleation days.

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Figure S10: Location of sampling site.

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Table 1: Ions identified by CI-APi-ToF

<u>Ion</u>	<u>m/Q</u>
<u>Cl⁻</u>	<u>34.97</u>
<u>NO₂⁻</u>	<u>45.99</u>
<u>C₃H₃O⁻</u>	<u>55.02</u>
<u>(NO₃)⁻</u>	<u>61.99</u>
<u>C₃H₃O₂⁻</u>	<u>71.01</u>
<u>C₃H₅O₂⁻</u>	<u>73.03</u>
<u>Br⁻</u>	<u>78.92</u>
<u>H₂O(NO₃)⁻</u>	<u>80.00</u>
<u>C₄H₅O₂⁻</u>	<u>85.03</u>
<u>C₃H₃O₃⁻</u>	<u>87.01</u>
<u>CH₃SO₃⁻</u>	<u>94.98</u>
<u>CFH₃(NO₃)⁻</u>	<u>96.01</u>
<u>HSO₄⁻</u>	<u>96.96</u>
<u>HCl(NO₃)⁻</u>	<u>97.97</u>
<u>H₄O₂(NO₃)⁻</u>	<u>98.01</u>
<u>C₄H₅O₃⁻</u>	<u>101.02</u>
<u>C₃H₃O₄⁻</u>	<u>103.00</u>
<u>SO₅⁻</u>	<u>111.95</u>
<u>C₄H₃O₄⁻</u>	<u>115.00</u>
<u>C₄H₅O₄⁻</u>	<u>117.02</u>
<u>C₃H₄O(NO₃)⁻</u>	<u>118.01</u>
<u>C₃H₃O₅⁻</u>	<u>119.00</u>
<u>HNO₃NO₃⁻</u>	<u>124.98</u>
<u>I⁻</u>	<u>126.91</u>
<u>C₄H₃O₅⁻</u>	<u>131.00</u>
<u>C₅H₇O₄⁻</u>	<u>131.03</u>
<u>C₄H₅O₅⁻</u>	<u>133.01</u>
<u>NH₃(HNO₃)(NO₃)⁻</u>	<u>142.01</u>
<u>C₅H₅O₅⁻</u>	<u>145.01</u>
<u>C₅H₇O₅⁻</u>	<u>147.03</u>
<u>C₃H₆O₃(NO₃)⁻</u>	<u>152.02</u>
<u>C₆H₇O₅⁻</u>	<u>159.03</u>
<u>C₄H₆O₃(NO₃)⁻</u>	<u>164.02</u>

<u>$C_3H_5NO_3(NO_3)^-$</u>	<u>165.02</u>
<u>$C_3H_4O_4(NO_3)^-$</u>	<u>166.00</u>
<u>$C_6H_5NO(NO_3)^-$</u>	<u>169.03</u>
<u>$C_2H_7N(HNO_3)(NO_3)^-$</u>	<u>170.04</u>
<u>$C_7H_7O_5^-$</u>	<u>171.03</u>
<u>$C_7H_9O_5^-$</u>	<u>173.05</u>
<u>IO_3^-</u>	<u>174.89</u>
<u>$C_4H_5NO_3(NO_3)^-$</u>	<u>177.02</u>
<u>$C_4H_4O_4(NO_3)^-$</u>	<u>178.00</u>
<u>$C_5H_7O_7^-$</u>	<u>179.02</u>
<u>$C_4H_6O_4(NO_3)^-$</u>	<u>180.01</u>
<u>$C_5H_9O_7^-$</u>	<u>181.04</u>
<u>$C_4H_8O_4(NO_3)^-$</u>	<u>182.03</u>
<u>$C_8H_{11}O_5^-$</u>	<u>187.06</u>
<u>$(HNO_3)_2(NO_3)^-$</u>	<u>187.98</u>
<u>$C_7H_9O_6^-$</u>	<u>189.04</u>
<u>$C_5H_7NO_3(NO_3)^-$</u>	<u>191.03</u>
<u>$C_4H_6N_2O_3(NO_3)^-$</u>	<u>192.03</u>
<u>$C_4H_5NO_4(NO_3)^-$</u>	<u>193.01</u>
<u>$C_5H_8O_4(NO_3)^-$</u>	<u>194.03</u>
<u>$H_2SO_4HSO_4^-$</u>	<u>194.93</u>
<u>$C_5H_7O_8^-$</u>	<u>195.01</u>
<u>$C_3H_6N_2O_4(NO_3)^-$</u>	<u>196.02</u>
<u>$C_4H_8O_5(NO_3)^-$</u>	<u>198.03</u>
<u>$C_4H_{11}N(HNO_3)(NO_3)^-$</u>	<u>198.07</u>
<u>$C_7H_7NO_2(NO_3)^-$</u>	<u>199.04</u>
<u>$C_6H_5NO_3(NO_3)^-$</u>	<u>201.02</u>
<u>$C_8H_{11}O_6^-$</u>	<u>203.06</u>
<u>$C_6H_6O_4(NO_3)^-$</u>	<u>204.01</u>
<u>$NH_3(HNO_3)_2(NO_3)^-$</u>	<u>205.01</u>
<u>$C_6H_8O_4(NO_3)^-$</u>	<u>206.03</u>
<u>$C_5H_7NO_4(NO_3)^-$</u>	<u>207.03</u>
<u>$C_4H_6N_2O_4(NO_3)^-$</u>	<u>208.02</u>
<u>$C_5H_8O_5(NO_3)^-$</u>	<u>210.03</u>
<u>$C_4H_7NO_5(NO_3)^-$</u>	<u>211.02</u>
<u>$C_8H_6O_3(NO_3)^-$</u>	<u>212.02</u>

<u>$C_3H_5NO_6(NO_3)^-$</u>	<u>213.00</u>
<u>$C_{10}H_{13}O_5^-$</u>	<u>213.08</u>
<u>$C_4H_8O_6(NO_3)^-$</u>	<u>214.02</u>
<u>$C_7H_7NO_3(NO_3)^-$</u>	<u>215.03</u>
<u>$C_7H_6O_4(NO_3)^-$</u>	<u>216.01</u>
<u>$C_7H_9NO_3(NO_3)^-$</u>	<u>217.05</u>
<u>$C_7H_8O_4(NO_3)^-$</u>	<u>218.03</u>
<u>$C_7H_{10}O_4(NO_3)^-$</u>	<u>220.05</u>
<u>$C_6H_9NO_4(NO_3)^-$</u>	<u>221.04</u>
<u>$C_5H_8N_2O_4(NO_3)^-$</u>	<u>222.04</u>
<u>$C_{10}H_7O_6^-$</u>	<u>223.02</u>
<u>$C_5H_8O_6(NO_3)^-$</u>	<u>226.02</u>
<u>$C_4H_7NO_6(NO_3)^-$</u>	<u>227.02</u>
<u>$C_4H_6O_7(NO_3)^-$</u>	<u>228.00</u>
<u>$C_8H_9NO_3(NO_3)^-$</u>	<u>229.05</u>
<u>$C_7H_7NO_4(NO_3)^-$</u>	<u>231.03</u>
<u>$C_2H_7N(HNO_3)_2(NO_3)^-$</u>	<u>233.04</u>
<u>$C_7H_{10}O_5(NO_3)^-$</u>	<u>236.04</u>
<u>$C_7H_{12}O_5(NO_3)^-$</u>	<u>238.06</u>
<u>$C_{10}H_7O_7^-$</u>	<u>239.02</u>
<u>$C_4H_6N_2O_6(NO_3)^-$</u>	<u>240.01</u>
<u>$C_5H_8O_7(NO_3)^-$</u>	<u>242.02</u>
<u>$C_5H_{11}NO_6(NO_3)^-$</u>	<u>243.05</u>
<u>$C_5H_{10}O_7(NO_3)^-$</u>	<u>244.03</u>
<u>$C_9H_{12}O_4(NO_3)^-$</u>	<u>246.06</u>
<u>$C_7H_7NO_5(NO_3)^-$</u>	<u>247.02</u>
<u>$C_8H_{10}O_5(NO_3)^-$</u>	<u>248.04</u>
<u>$C_7H_9NO_5(NO_3)^-$</u>	<u>249.04</u>
<u>$C_8H_{12}O_5(NO_3)^-$</u>	<u>250.06</u>
<u>$C_7H_{11}NO_5(NO_3)^-$</u>	<u>251.05</u>
<u>$C_7H_{10}O_6(NO_3)^-$</u>	<u>252.04</u>
<u>$C_7H_{12}O_6(NO_3)^-$</u>	<u>254.05</u>
<u>$C_6H_{11}NO_6(NO_3)^-$</u>	<u>255.05</u>
<u>$C_6H_{10}O_7(NO_3)^-$</u>	<u>256.03</u>
<u>$C_5H_8O_8(NO_3)^-$</u>	<u>258.01</u>
<u>$C_8H_7NO_5(NO_3)^-$</u>	<u>259.02</u>

<u>$C_9H_{10}O_5(NO_3)^-$</u>	<u>260.04</u>
<u>$C_4H_{11}N(HNO_3)_2(NO_3)^-$</u>	<u>261.07</u>
<u>$C_9H_{12}O_5(NO_3)^-$</u>	<u>262.06</u>
<u>$C_7H_7NO_6(NO_3)^-$</u>	<u>263.02</u>
<u>$C_8H_{11}NO_5(NO_3)^-$</u>	<u>263.05</u>
<u>$C_8H_{10}O_6(NO_3)^-$</u>	<u>264.04</u>
<u>$C_9H_{14}O_5(NO_3)^-$</u>	<u>264.07</u>
<u>$C_7H_9NO_6^-(NO_3)^-$</u>	<u>265.03</u>
<u>$C_7H_8O_7(NO_3)^-$</u>	<u>266.01</u>
<u>$C_8H_{12}O_6(NO_3)^-$</u>	<u>266.05</u>
<u>$C_7H_{11}NO_6(NO_3)^-$</u>	<u>267.05</u>
<u>$C_8H_{15}NO_5(NO_3)^-$</u>	<u>267.08</u>
<u>$C_7H_{10}O_7(NO_3)^-$</u>	<u>268.03</u>
<u>$C_6H_{11}NO_7(NO_3)^-$</u>	<u>271.04</u>
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<u>$C_5H_9NO_8(NO_3)^-$</u>	<u>273.02</u>
<u>$C_5H_8O_9(NO_3)^-$</u>	<u>274.01</u>
<u>$C_{10}H_{12}O_5(NO_3)^-$</u>	<u>274.06</u>
<u>$C_{10}H_{14}O_5(NO_3)^-$</u>	<u>276.07</u>
<u>$C_9H_{13}NO_5(NO_3)^-$</u>	<u>277.07</u>
<u>$C_9H_{12}O_6(NO_3)^-$</u>	<u>278.05</u>
<u>$C_{10}H_{16}O_5(NO_3)^-$</u>	<u>278.09</u>
<u>$C_8H_{11}NO_6^-(NO_3)^-$</u>	<u>279.05</u>
<u>$C_9H_{14}O_6(NO_3)^-$</u>	<u>280.07</u>
<u>$C_7H_9NO_7(NO_3)^-$</u>	<u>281.03</u>
<u>$C_8H_{12}O_7(NO_3)^-$</u>	<u>282.05</u>
<u>$C_7H_{11}NO_7(NO_3)^-$</u>	<u>283.04</u>
<u>$C_8H_{14}O_7(NO_3)^-$</u>	<u>284.06</u>
<u>$C_{10}H_9NO_5(NO_3)^-$</u>	<u>285.04</u>
<u>$C_5H_8N_2O_8(NO_3)^-$</u>	<u>286.02</u>
<u>$C_5H_7NO_9(NO_3)^-$</u>	<u>287.00</u>
<u>$C_7H_{15}NO_7(NO_3)^-$</u>	<u>287.07</u>
<u>$C_5H_{10}N_2O_8(NO_3)^-$</u>	<u>288.03</u>
<u>$C_5H_9NO_9(NO_3)^-$</u>	<u>289.02</u>
<u>$C_{10}H_{12}O_6(NO_3)^-$</u>	<u>290.05</u>
<u>$C_{10}H_{15}NO_5(NO_3)^-$</u>	<u>291.08</u>

<u>C₁₀H₁₄O₆(NO₃)⁻</u>	<u>292.07</u>
<u>C₉H₁₃NO₆(NO₃)⁻</u>	<u>293.06</u>
<u>C₁₀H₁₆O₆(NO₃)⁻</u>	<u>294.08</u>
<u>C₈H₁₁NO₇(NO₃)⁻</u>	<u>295.04</u>
<u>C₉H₁₄O₇(NO₃)⁻</u>	<u>296.06</u>
<u>C₈H₁₃NO₇(NO₃)⁻</u>	<u>297.06</u>
<u>C₈H₁₂O₈(NO₃)⁻</u>	<u>298.04</u>
<u>C₇H₁₀O₉(NO₃)⁻</u>	<u>300.02</u>
<u>C₁₀H₉NO₆(NO₃)⁻</u>	<u>301.03</u>
<u>C₇H₁₂O₉(NO₃)⁻</u>	<u>302.04</u>
<u>C₁₀H₁₁NO₆(NO₃)⁻</u>	<u>303.05</u>
<u>C₆H₁₀O₁₀(NO₃)⁻</u>	<u>304.02</u>
<u>C₁₂H₁₈O₅(NO₃)⁻</u>	<u>304.10</u>
<u>C₉H₉NO₇(NO₃)⁻</u>	<u>305.03</u>
<u>C₉H₈O₈(NO₃)⁻</u>	<u>306.01</u>
<u>C₁₁H₁₆O₆(NO₃)⁻</u>	<u>306.08</u>
<u>C₁₀H₁₅NO₆(NO₃)⁻</u>	<u>307.08</u>
<u>C₁₀H₁₄O₇(NO₃)⁻</u>	<u>308.06</u>
<u>C₉H₁₃NO₇(NO₃)⁻</u>	<u>309.06</u>
<u>C₁₀H₁₆O₇(NO₃)⁻</u>	<u>310.08</u>
<u>C₉H₁₅NO₇(NO₃)⁻</u>	<u>311.07</u>
<u>C₉H₁₄O₈(NO₃)⁻</u>	<u>312.06</u>
<u>C₈H₁₃NO₈(NO₃)⁻</u>	<u>313.05</u>
<u>C₈H₁₂O₉(NO₃)⁻</u>	<u>314.04</u>
<u>C₇H₁₁NO₉(NO₃)⁻</u>	<u>315.03</u>
<u>C₁₀H₁₀N₂O₆(NO₃)⁻</u>	<u>316.04</u>
<u>C₁₀H₉NO₇(NO₃)⁻</u>	<u>317.03</u>
<u>C₁₂H₁₆O₆(NO₃)⁻</u>	<u>318.08</u>
<u>C₁₀H₁₁NO₇(NO₃)⁻</u>	<u>319.04</u>
<u>C₁₀H₁₀O₈(NO₃)⁻</u>	<u>320.03</u>
<u>C₁₂H₁₈O₆(NO₃)⁻</u>	<u>320.10</u>
<u>C₁₀H₁₃NO₇(NO₃)⁻</u>	<u>321.06</u>
<u>C₁₁H₁₆O₇(NO₃)⁻</u>	<u>322.08</u>
<u>C₁₀H₁₅NO₇(NO₃)⁻</u>	<u>323.07</u>
<u>C₁₀H₁₄O₈(NO₃)⁻</u>	<u>324.06</u>
<u>C₁₀H₁₆O₈(NO₃)⁻</u>	<u>326.07</u>

<u>C₉H₁₅NO₈(NO₃)⁻</u>	<u>327.07</u>
<u>C₉H₁₄O₉(NO₃)⁻</u>	<u>328.05</u>
<u>C₈H₁₃NO₉(NO₃)⁻</u>	<u>329.05</u>
<u>C₉H₁₆O₉(NO₃)⁻</u>	<u>330.07</u>
<u>C₁₁H₁₁NO₇(NO₃)⁻</u>	<u>331.04</u>
<u>C₁₀H₉NO₈(NO₃)⁻</u>	<u>333.02</u>
<u>C₁₂H₁₆O₇(NO₃)⁻</u>	<u>334.08</u>
<u>C₁₂H₁₈O₇(NO₃)⁻</u>	<u>336.09</u>
<u>C₁₀H₁₃NO₈(NO₃)⁻</u>	<u>337.05</u>
<u>C₁₁H₁₆O₈(NO₃)⁻</u>	<u>338.07</u>
<u>C₁₀H₁₅NO₈(NO₃)⁻</u>	<u>339.07</u>
<u>C₁₀H₁₄O₉(NO₃)⁻</u>	<u>340.05</u>
<u>C₁₃H₁₃NO₆(NO₃)⁻</u>	<u>341.06</u>
<u>C₁₀H₁₆O₉(NO₃)⁻</u>	<u>342.07</u>
<u>C₉H₁₅NO₉(NO₃)⁻</u>	<u>343.06</u>
<u>C₉H₁₄O₁₀(NO₃)⁻</u>	<u>344.05</u>
<u>C₁₂H₁₃NO₇(NO₃)⁻</u>	<u>345.06</u>
<u>C₁₃H₁₆O₇(NO₃)⁻</u>	<u>346.08</u>
<u>C₁₃H₁₈O₇(NO₃)⁻</u>	<u>348.09</u>
<u>C₁₂H₁₇NO₇(NO₃)⁻</u>	<u>349.09</u>
<u>C₁₃H₂₀O₇(NO₃)⁻</u>	<u>350.11</u>
<u>C₁₁H₁₅NO₈(NO₃)⁻</u>	<u>351.07</u>
<u>C₁₁H₁₇NO₈(NO₃)⁻</u>	<u>353.08</u>
<u>C₁₀H₁₅NO₉(NO₃)⁻</u>	<u>355.06</u>
<u>C₁₁H₁₈O₉(NO₃)⁻</u>	<u>356.08</u>
<u>C₁₀H₁₇NO₉(NO₃)⁻</u>	<u>357.08</u>
<u>C₁₃H₁₄O₈(NO₃)⁻</u>	<u>360.06</u>
<u>C₁₂H₁₃NO₈(NO₃)⁻</u>	<u>361.05</u>
<u>C₁₃H₁₉NO₇(NO₃)⁻</u>	<u>363.10</u>
<u>C₁₃H₁₈O₈(NO₃)⁻</u>	<u>364.09</u>
<u>C₁₂H₁₇NO₈(NO₃)⁻</u>	<u>365.08</u>
<u>C₁₃H₂₀O₈(NO₃)⁻</u>	<u>366.10</u>
<u>C₁₂H₁₉NO₈(NO₃)⁻</u>	<u>367.10</u>
<u>C₁₂H₁₈O₉(NO₃)⁻</u>	<u>368.08</u>
<u>C₁₁H₁₇NO₉(NO₃)⁻</u>	<u>369.08</u>
<u>C₁₁H₁₆O₁₀(NO₃)⁻</u>	<u>370.06</u>

<u>$C_{10}H_{15}NO_{10}(NO_3)^-$</u>	<u>371.06</u>
<u>$C_{14}H_{14}O_8(NO_3)^-$</u>	<u>372.06</u>
<u>$C_{13}H_{13}NO_8(NO_3)^-$</u>	<u>373.05</u>
<u>$C_{14}H_{16}O_8(NO_3)^-$</u>	<u>374.07</u>
<u>$C_{14}H_{20}O_8(NO_3)^-$</u>	<u>378.10</u>
<u>$C_{13}H_{19}NO_8(NO_3)^-$</u>	<u>379.10</u>
<u>$C_{14}H_{22}O_8(NO_3)^-$</u>	<u>380.12</u>
<u>$C_{13}H_{21}NO_8(NO_3)^-$</u>	<u>381.12</u>
<u>$C_{13}H_{20}O_9(NO_3)^-$</u>	<u>382.10</u>
<u>$C_{16}H_{19}NO_6(NO_3)^-$</u>	<u>383.11</u>
<u>$C_{16}H_{18}O_7(NO_3)^-$</u>	<u>384.09</u>
<u>$C_{15}H_{17}NO_7(NO_3)^-$</u>	<u>385.09</u>
<u>$C_{15}H_{16}O_8(NO_3)^-$</u>	<u>386.07</u>
<u>$C_{10}H_{15}NO_{11}(NO_3)^-$</u>	<u>387.05</u>
<u>$C_{11}H_{18}O_{11}(NO_3)^-$</u>	<u>388.07</u>
<u>$C_{18}H_{21}N(NO_3)O_5^-$</u>	<u>393.13</u>
<u>$C_{15}H_{24}(NO_3)O_8^-$</u>	<u>394.14</u>
<u>$C_{18}H_{23}N(NO_3)O_5^-$</u>	<u>395.15</u>
<u>$C_{18}H_{22}(NO_3)O_6^-$</u>	<u>396.13</u>
<u>$C_{17}H_{21}N(NO_3)O_6^-$</u>	<u>397.13</u>
<u>$C_{17}H_{20}(NO_3)O_7^-$</u>	<u>398.11</u>
<u>$C_{16}H_{19}N(NO_3)O_7^-$</u>	<u>399.10</u>
<u>$C_{16}H_{18}(NO_3)O_8^-$</u>	<u>400.09</u>
<u>$C_{15}H_{17}N(NO_3)O_8^-$</u>	<u>401.08</u>
<u>$C_{12}H_{20}(NO_3)O_{11}^-$</u>	<u>402.09</u>
<u>$C_{10}H_{15}N(NO_3)O_{12}^-$</u>	<u>403.05</u>
<u>$C_{15}H_{18}(NO_3)O_9^-$</u>	<u>404.08</u>
<u>$C_{18}H_{17}N(NO_3)O_6^-$</u>	<u>405.09</u>
<u>$C_{19}H_{23}N(NO_3)O_5^-$</u>	<u>407.15</u>
<u>$C_{19}H_{22}(NO_3)O_6^-$</u>	<u>408.13</u>

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