

## Response to reviewer #1

We thank reviewer #1 for the positive assessment of our work. Reviewer comments are in italics.

*I have enjoyed reading the manuscript and find it a remarkable piece of work, stretching from individual measurements to continental scale impact estimates. At each level, interesting results are presented, starting with uncertainties connected with assumptions regarding emission potential and temperature sensitivity in standard emission models. The analysis is timely and thorough although the literature overview could be a bit more comprehensive. The text is well written and the conclusions are sound. My only concern is that disregarding the CO<sub>2</sub> effect in the scenario analyses might call for exaggerated calls for action. Thus, I would appreciate a 7th simulation to account for this, e.g. a scenario such as the 6th run but with increased CO<sub>2</sub>, despite the fact that I am aware of the inconsistency with the measurements (and also with some uncertainties C1 related to a shift of temperature response curves under higher CO<sub>2</sub> as shown by Sun et al.).*

The CO<sub>2</sub> effect is easy to add, and we have included a 7<sup>th</sup> simulation as suggested.

The text at line 239 has been changed to say: "The climate2050 run does not include the associated increases in CO<sub>2</sub> mixing ratios, to be consistent with our measurements which were also not conducted in a higher CO<sub>2</sub> atmosphere. A 7th simulation assumes a 550 ppm CO<sub>2</sub> atmosphere on top of the delta-scaled surface temperatures, employing Heald et al's (2009) method for calculating short and long term CO<sub>2</sub> activity factors,  $\gamma$ C. Fixing the atmospheric CO<sub>2</sub> to 550 ppm reduces the isoprene emissions by 5% in the short term and 13% in the long term."

An extra column has been added to table 4 to show simulation 7 has  $\gamma$ C added.

Add results from this experiment to line 280 "The addition of a higher CO<sub>2</sub> atmosphere has reduced the daytime isoprene by 15 - 26 % from the climate2050 run, across the three campaigns."

To avoid too many panels in figure 5, we'll leave the climate 2050 run as the upper end to the range in results. However, figure 6 now contains the time series from the climate2050\_ $\gamma$ C run. Figure 6 has been altered to have the Sydney results on the left and the Darwin results on the right.

Change text at line 310 "However peak O<sub>3</sub> in Sydney increases by 10 – 15 ppb as an hourly average in the FC\_ $\gamma$ T+LEF differences, but by 12 - 17 ppb in the climate2050\_ $\gamma$ C differences and 15 - 21 ppb in the climate2050 differences (Figure 6a,b). These increases represent 10 - 21 % of the O<sub>3</sub> NEPM."

Change the text at line 320 "The climate2050 (and  $\gamma$ C) differences show days with an increase of 0.42  $\mu$ g m<sup>-3</sup> in Sydney and 0.14  $\mu$ g m<sup>-3</sup> in Darwin (2 % and 1 % of the PM<sub>2.5</sub> 2025 NEPM, respectively)."

Make minor changes to conclusions at line 350 "Three future experiments were conducted, the first using current day meteorology, the second using a delta-scaled surface temperature change to projected 2050 summertime temperatures, and the third using a 550 ppm atmospheric CO<sub>2</sub> on top of the delta scaled temperatures."

And at line 354. "The climate2050 experiment showed much larger increases in isoprene, O<sub>3</sub> and biogenic SOA across Australia, tempered slightly by the addition of increased atmospheric CO<sub>2</sub>."

Also adjust abstract at line 13 to include "A 550 ppm CO<sub>2</sub> atmosphere in 2050 mitigates these peak Sydney O<sub>3</sub> mixing ratios by 4 ppb. Nevertheless, these forecasted increases in O<sub>3</sub> are up to one fifth..."

*From the few specific remarks, I would like to stress the benefits from an improved literature overview (i.e. L25ff). The generally high emission potential of eucalypts in comparison with other species have firstly been depicted in Evans et al. 1982 and can also be derived from Kesselmeier and Staudt – although the values concentrate on E.globulus. Karlik and Winer as well as Geron et al. provide an additional emission rate of E.camaldulensis (28, 14.6, add to Table 1) and a couple of other eucalypt species – although not under Australian conditions.*

Add text at line 26 “Native to Australia, eucalypt trees are amongst the highest BVOC emitters of any plant species (Benjamin et al., 1996; Evans et al., 1982; Kesselmeier and Staudt, 1999)”

Add *E. camaldulensis* emission rates from Karlik and Winer (2001) to table 1. The value listed in the Geron et al (2001) reference is the measurement by He et al (2000) already listed in table 1. Geron et al convert He et al’s measurement to a dry weight of carbon ( $\mu\text{g C g}^{-1} \text{h}^{-1}$ ) which is why the value is slightly different. Here we’re using  $\mu\text{g}$  isoprene  $\text{g}^{-1} \text{h}^{-1}$ .

*You may also note that different temperature responses and emission factor variability have been obtained before, starting with the original Guenther et al. 1991 publications and widely discussed e.g. in Niinemets et al. 2010 (e.g. L44ff).*

True, and I noted from Guenther et al (1993) when one of the early models for isoprene emission with temperature was defined, it was based on “empirical coefficients which were determined by nonlinear best fit procedures using eucalyptus, sweet gum, aspen, and velvet bean emission rate measurements.”

Change text at line 45 “Whilst the MEGAN parameterisations are fitted from a wide range of ecosystem responses to environmental conditions, there are spatial and temporal exceptions to these standards which are comprehensively reviewed by Niinemets et al (2010)”

*In the end of chapter 2.2 (93ff), I got the impression that the authors are carried away a bit. First, the last paragraph seems to fit better into a discussion; and second, the first sentence is not logical (the measurements are hardly going to change but emission rates and species abundance probably will). By the way, I am still uncertain to which degree these 4 eucalypt species are actually representative for the Australian forests or how abundant they are in relative terms (L163).*

We’ve decided to drop the first part of this paragraph and include more details of how the eucalypt species are spread in the earlier part of section 2.2. We also include species occurrence maps from the Atlas of Living Australia in the supplementary section, reproduced below.

At line 75 “*E. camaldulensis* and *E. tereticornis* have a wide geographical representation within Australia, with a latitudinal native growing range of 9-38 °S (Atlas of Living Australia, 2019), (supplementary figure S1). *E. camaldulensis* is the most widely naturally distributed species of all eucalypts in Australia (Atlas of Living Australia, 2019). The native climatic distribution range of *E. botryoides* and *E. smithii* are restricted to the south east coastal regions. All four species are forecast to exist in future, but only *E. camaldulensis* is predicted to expand its growing area by 2085 (González-Orozco et al., 2016).”

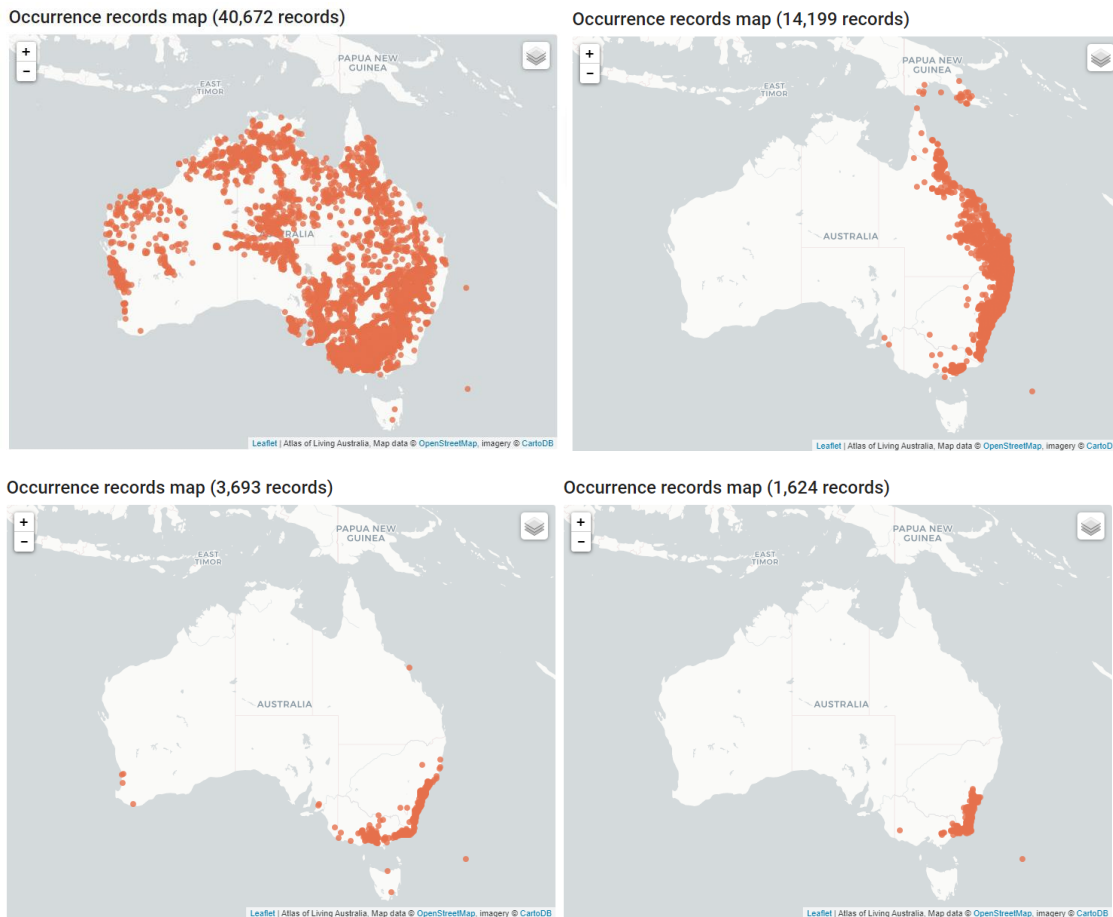


Figure S1 clockwise from top left. *E. camaldulensis*, *E. tereticornis*, *E. smithii*, *E. botryoides*.

*With one sole reference, the protective functions of isoprene to sun flecks and very high temperatures are not very well acknowledged (L268ff). There are several publications (e.g. Behnke et al.) and reviews (Loreto and Fineshi) that illustrate this function. In fact, emission is prolonged even under carbon deficit conditions (Yañez-Serrano et al.).*

At line 270 “Hot and windy conditions would cause lots of sun-flecking within the tree canopy, causing sudden temperature spikes on the leaf surface. Physiologically, the increased production of isoprene during temperature and light spikes helps to maintain photosynthesis (Behnke et al., 2010) during times of mild stresses (Loreto and Fineschi, 2015), above and beyond leaf cooling via transpiration processes (Sharkey et al., 2008). High isoprene emitters can better survive prolonged heatwaves (Yañez-Serrano et al., 2019), although Aspinwall et al’s (2019) study on our four eucalypt species showed trees grown under future climate conditions suffered greater heatwave damage than the same species in current climate conditions.”

*Finally, the thought came to my mind that instead of removing the trees (which is of course not recommended), the forest management might be compelled to introduce species others than Eucalypts that are not emitting isoprene (L314ff). However, given the protective function mentioned above, this option might not be advisable because non-emitters might not be able to withstand the coming heat (Penuelas and Munne-Bosch, Ryan et al.).*

We’re not recommending eucalypt trees are removed, though will add a sentence about non emitters being unable to cope with heat stresses.

At line 315 that “new urban developments should consider the BVOC emission potential of trees before planting (Paton-Walsh et al., 2019), taking into account that non or low emitting trees may not withstand climate induced heatwaves (Peñuelas and Munné-Bosch, 2005).”

## **Response to reviewer #2**

We thank Reviewer #2, who added their comments in the form of an annotated PDF document. We hope we caught them all below! Reviewer comments are in italics.

*Discussion on use of saplings in the discussion. I think this needs to come back even more in the discussion - what impact then would this have on your findings/conclusions?*

We have added discussion at line 347 (refer to penultimate comment)

*Picture of the leaf cuvette with further details on experiment. A picture of this might help as I am not familiar with it. Does the cuvette cover the full leaf? Also, how long was the cuvette exposed? Was this done in-line with PTR-MS? So continuously? Or every 7-minutes?*

We feel a picture of a commercial leaf cuvette doesn't add to the manuscript. However, we have added more details to explain that gas exchange measurements were continuous and that the full leaf was measured. Leaf area measurements are already described in L108-9 of the original submission.

At line 98: “Leaf gas exchange measurements were made continuously with a LI-6400XT portable photosynthesis system (Li-Cor Inc., Lincoln, NE, USA) connected to a Walz 3010-GWK1 leaf cuvette (maximum surface area for leaf 140 cm<sup>2</sup>; Heinz Walz GmbH, Effeltrich, Germany).”

*Were the results of these 5-6 trees quite similar? So we can assume this number of points is representative of each species?*

The values obtained from the 5-6 replicated were similar even though there was intrinsic variability within a species. To demonstrate this we have added a figure to the supplementary section that highlights the trends in response quite clearly.

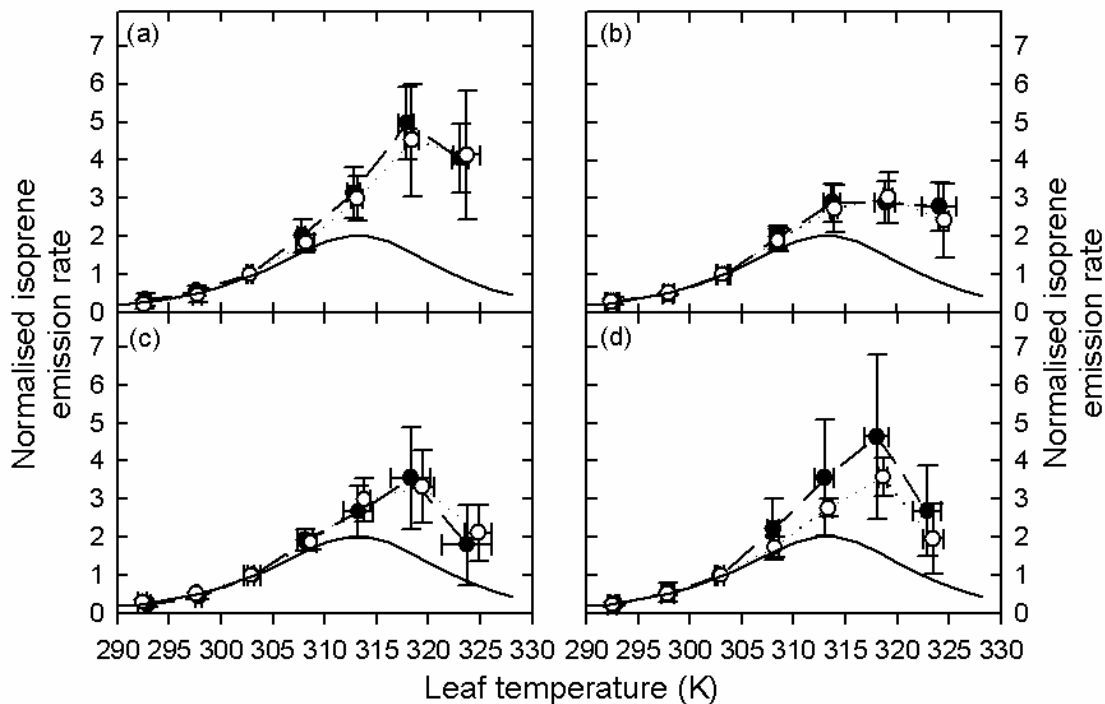


Figure: Temperature response of normalised isoprene emission rate from four *Eucalyptus* species (a) *E. camaldulensis*, (b) *E. botryoides*, (c) *E. smithii* and (d) *E. tereticornis* grown under two different temperature regimes. Open circles (dotted lines) are current climate and filled circles (dashed line) are future climate. The solid line in each panel is the normalised isoprene emission calculated using default MEGAN values. Data are normalised to the isoprene emission rate measured at a leaf temperature of 303 K. Error bars (horizontal and vertical) are means  $\pm$  one standard deviation of 4-6 replicate plants.

*It is not clear at this point in the text of the importance of 303K and 313K. I would recommend stating the importance of these thresholds here.*

These values relate to perceived thresholds at 30C and 40C. The point of this sentence is to convey how hot these three summers were, so will change to maximum and averages over the campaigns. Alter text at line 131 "Maximum (and average) measured temperatures were 308.9 K (295.9 K) for Bringelly, 310.0 K (295.6 K) for SPS1 and 317.2K (295.3 K) for MUMBA. Climate projections for Australia forecast increases in average temperatures with an accompanying increase in the frequency of extreme heatwave days (Bureau of Meteorology and CSIRO, 2018)."

Also remove text at line 243. Move "Delta-scaling adds  $\sim$ 2 K to the surface temperatures near Sydney" to line 233.

*If your emissions are lower than He who used a mixture of leaf ages, could your results be a partial artefact of not having old leaves?*

Our basal emission rates were higher than He's and are consistent with the observations of Street (1997) that younger leaves are higher emitters.

Re-write from line 193: “He et al. (2000) used a mixture of young and mature leaves in their experiments, which could be one explanation for the difference in emission rates as young leaves are expected to be higher emitters than older leaves in *Eucalyptus* (Street et al. 1997). However, as the growth conditions (particularly light and temperature) and measurement protocols between this study and He et al. (2000) were different (we directly measured BER with a leaf cuvette at  $1000 \mu\text{mol m}^{-2} \text{s}^{-1}$  and  $30 \text{ }^\circ\text{C}$  while He et al. used a dynamic chamber and scaled emissions to  $1000 \mu\text{mol m}^{-2} \text{s}^{-1}$  PAR and  $30 \text{ }^\circ\text{C}$  using algorithms from Guenther et al. (1993)), it is difficult to undertake a direct comparison. However, our measurements put the four eucalypt species into the high emission category.”

*Are these averages for 4 species in MEGAN or in your data or both? I assume it can't be both as there wasn't info on all your species before. So then Megan is for the three species?*

The values given for MEGAN in table 3 is the average across the isoprene emission factor map for the inner 3km domain. This is stated in the table caption. On line 31 we state that the eucalypt tree emissions in MEGAN were based on six studies with numerous species involved. The default isoprene emission factor map represents all tree, shrub and grass species in Australia (line 201). What we have done is then alter the tree portion only of this map using our 4 tree species measurements but weighted them according to the area each species takes up in Australia (line 208).

To make this clearer we alter text at line 207 “Table 3 shows how the results of fitting CT1, CT2, Tmax and Ceo compared to the default MEGAN values. These new fitted data are for the four tree species in the experiment, weighted according to their coverage in Table 1. The new average LEFs from our four eucalypt species are 31-48 % lower than the default average MEGAN LEF we use in the base run for the 3km Sydney domain.”

*What does the 40% emission reduction relate to? estimated here for what? I would recommend giving more context to this statement.*

At line 210: “Previous modelling showed that a 40% reduction in isoprene was needed to better match the observations from our three field campaigns (Emmerson et al., 2019).”

*New numbers to default MEGAN? What are the two levels used for each variable? What do you change them to in this sensitivity? Is it MEGAN vs new numbers?*

The two levels in table 3 refer to the measurements made on the current climate and future climate grown trees as described in section 2.2. The sensitivity involves changing the default MEGAN values to the two values in columns 3 and 4 of table 3.

Alter text at line 211 “The value fitted for CT<sub>2</sub> is very high ( $1158.36 \text{ kJ mol}^{-1}$ ) in the future climate treatment compared with the current climate treatment ( $167.11 \text{ kJ mol}^{-1}$ ) and default MEGAN ( $230 \text{ kJ mol}^{-1}$ ), due to the future climate *E. camaldulensis* measurements in Figure 2.”

*Thus, CT2 won't be re-fitted, correct?*

Correct.

At line 217 “The high CT<sub>2</sub> value in the future climate treatment will not be refitted, as the incurred 19 % decrease in isoprene is small compared with the 282 % increase caused by C<sub>eo</sub>.”

*(relates to figure 3) what are the LEFs normalised to? Assume that these are max emission points in summer? what are the emissions normalised to? Are these the mean points? Is there deviations as this is across multiple field campaigns? As these are all Jan-Mar, I assume that is max emission period, correct?*

The LEFS do not change regardless of season. Instead they are moderated in MEGAN by the environmental factors mentioned on line 44 which accounts for seasonal differences (temperature, PAR, leaf area index, leaf age, soil moisture, and suppression via ambient CO<sub>2</sub> concentrations). The LEFS in this figure are normalised by the default MEGAN LEF from table 3. I.e., default MEGAN LEF = 1, CC\_γT+LEF is 48% less and FC\_γT+LEF is 31% less than the default.

Change the description of the figure at line 247 “If the leaf temperature is varied within Equations 1-4 and γT is multiplied by the LEF, the impacts of experiments 1-5 on isoprene emission start at about 283 K (Figure 3). Experiment 6 follows the FC\_γT+LEF profile. Here, the new current and future climate LEFs are normalised by the default MEGAN LEF.”

*due to spatial heterogeneity and proximity to study site, right? The next sentence is a little confusing of an explanation to me, so if possible I would recommend summarizing it here before explanation*

At line 256: “While it is intuitive to expect less isoprene will be emitted in the CC\_γT+LEF and FC\_γT+LEF experiments over the base run (from Figure 3), this may not be the case due to spatial heterogeneity in the new current and future climate LEF maps.”

*Where are these decreases on Fig 3? I don't see them.*

Change text at line 259: “The results from experiments 3 and 5 certainly show a sustained isoprene decrease below 314 K and 311 K respectively.”

*The colour between Obs and Base hard to see in legend. Is it possible to show also circle with line in "observation".*

Done

*Are these field campaigns at one site for each campaign? So are modelled output extracted for that one point? Or full domain?*

At line 263: “The C-CTM is compiled with changes to MEGAN implemented according to Table 3, run for experiments 1-6 (Table 4) and the isoprene time series is extracted at each field campaign site. The modelled mean diurnal profiles of isoprene are then compared to the mean diurnal observations taken at each field campaign (Figure 4).”

*As seen in fig 3?*

Correct. At line 265 “The CC\_γT variables only increase the isoprene mixing ratios when temperatures exceed 303 K (from figure 3)”

How were  $r^2$  calculated? Is this based on all hourly values for each field campaign? Were there similar number of points? Also, why are there missing points in the observations? Did you use a data completeness threshold or something to develop these?

The gaps in the observations are at specific times for blanks/calibrations, eg 2 and 3am local time. The  $r^2$  were calculated by comparing the mean diurnal modelled isoprene to the mean diurnal observed isoprene. However the  $r^2$  is similar if all hourly data is included.

At line 265 "Instrument calibrations/blanks are taken at least twice a day, incurring frequent regular gaps in observed isoprene."

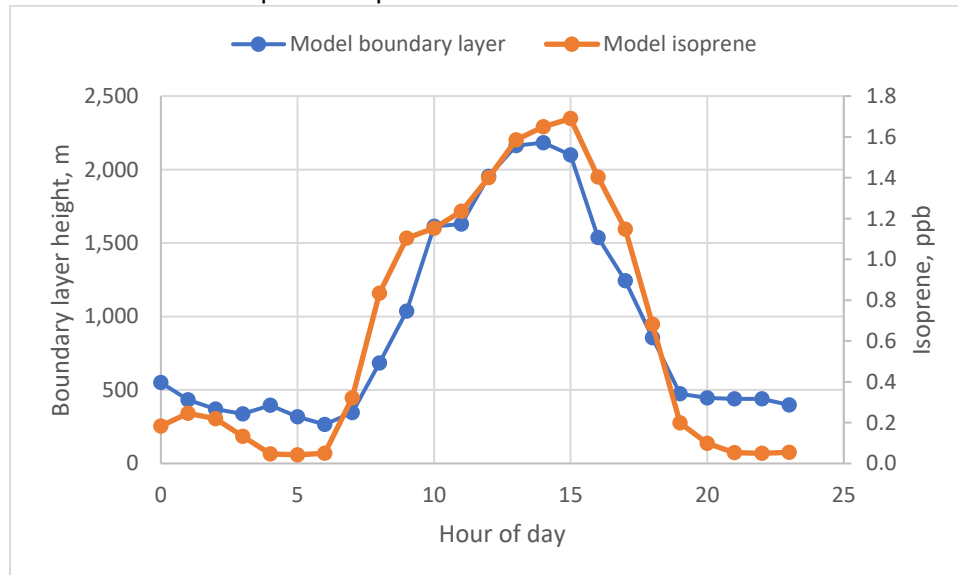
At line 266 "but generally the CC\_γT and CC\_γT+LEF experiments have increased the diurnal modelled to observed  $r^2$  when compared with the  $r^2$  between the base run and observations."

Comment refers to continuing large bias in MUMBA results.

At line 276 "The CC\_γT+LEF experiments represent current day conditions, with roughly the correct magnitude (MUMBA excepted) of predicted isoprene and best statistical fit compared with the observations."

Refers to nighttime decrease in isoprene. It is still higher than measurements in MUMBA - and then is too small 4-8:00 in SPS1. I would recommend adding some more discussion on the biases in this section. In addition, it might be helpful to show in supp materials only the best performing one(s) (cc\_T+LEF) with base and observations with std deviations - to show also the spread of the data and if these are within each other by 1 std dev or not.

The biases in pre-dawn SPS1 are because there is a very slight rise in the boundary layer, causing dilution of the atmospheric isoprene:



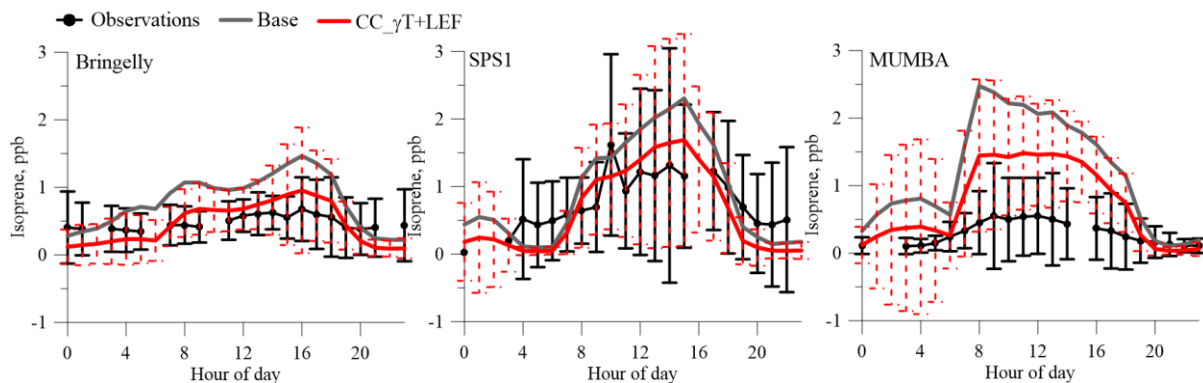
At line 286 "Conversely a slight rise in the model boundary layer at 04:00 AEDT in SPS1 causes dilution of the atmospheric isoprene."

We have always been aware of high modelled biases in Australian isoprene modelling, particularly during MUMBA (the lead author has three other publications on the isoprene bias, cited in this manuscript). We have made the std deviation plot as requested (below), although the colour of the CC\_γT+LEF run was changed to red to show up better. There is wide variation in observed isoprene, particularly during SPS1. The model also shows high variation before dawn in SPS1 and MUMBA. We



calculate the percentage of modelled hours which are within +1 standard deviation of the observed isoprene as follows:

Base run (CC\_γT+LEF run). Bringelly = 40% (90%); SPS1 = 89% (100%); MUMBA = 19% (33%).



The point of this paper was to see if we could use the new measurements to reduce the modelled biases through the sensitivity runs, which we have done. The impacts of the future simulations are then explored via their differences on the least biased current climate run, not absolute values. I think we've made the point that the modelled isoprene was never perfect to start with, and much additional discussion of the biases is not necessary. We will include the figure in the supplementary.

At line 267 "The average modelled isoprene in the CC\_γT+LEF run is within +/- 1 standard deviation of the observations 90 – 100 % of the time during Bringelly and SPS1, and 33 % during MUMBA which continues to exhibit high bias (supplementary figure S4)".

*In which year was air quality in Sydney and Darwin classed as 'very good' 2019? Or in many years?*  
This relates to the latest Australian State of the Environment reporting period of 2009-2014.  
Line 309 "(years 2009 – 2014)"

*Improving the LEF did more to decrease the bias than the temperature response measurements? I think this is an important findings. Looking at the model vs observations, improving the LEF did more to decrease bias than the improved temperature response - i.e. dark blue line compared obs vs light blue line compare to obs.*

We knew that decreasing the LEF by about 40% would bring the magnitude of the model into better agreement with the observations from our previous work (see above). However, the largest increase in the  $r^2$  fit of the data actually occurred when using the temperature response data alone (for Bringelly and MUMBA) than the decrease in LEF. Correctly fitting the SPS1 observed data is more difficult due to the spike at 10am.

*Would peak in temperature response be different if the measurements were conducted on older leaves? Would it be expected that the peak temp in the temp response also be different? Or is it just the total emissions? As I mentioned earlier, I would recommend that this issue about new and old leaves and the impact here is unpacked a bit more.*

Without further experimentation using older trees it is impossible to answer the reviewer's question and any comment would be pure speculation. However, in relation to the reviewer's earlier comment, we have added some more discussion at line 347.

“Our measurements were conducted on sapling trees which may exhibit higher isoprene emissions than adult trees when emission rates are expressed on leaf mass basis but not on a leaf area basis (Street et al., 1997). Street et al. (1997) explained this through younger leaves having a higher specific leaf area than older leaves because eucalypts exhibit heterophylly (the foliage leaves on the same plant are of two distinctly different types). The apparent difference in emission rates between young and old leaves could be a consequence of morphology rather than biochemistry, so we expect the trend between the current and future climate growth emissions to be similar amongst trees of all ages.”

*Similar to a comment above - which eucalyptus were used in MEGAN to estimate these characteristics?*

See comment above

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# Temperature response measurements from eucalypts give insight into the impact of Australian isoprene emissions on air quality in 2050

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**Abstract.** Predicting future air quality in Australian cities dominated by eucalypt emissions requires an understanding of their emission potentials in a warmer climate. Here we measure the temperature response in isoprene emissions from saplings of four different *Eucalyptus* species grown under current and future average summertime temperature conditions. The future conditions represent a 2050 climate under Representative Concentration Pathway 8.5, with average daytime temperatures of 294.5 K. Ramping the temperature from 293 K to 328 K resulted in these eucalypts emitting isoprene at temperatures 4-9 K higher than default maximum emission temperature in the Model of Emissions of Gases and Aerosols from Nature (MEGAN). New basal emission rate measurements were obtained at the standard conditions of 303 K leaf temperature and 1000  $\mu\text{mol m}^{-2} \text{s}^{-1}$  photosynthetically active radiation and converted into landscape emission factors. We applied the eucalypt temperature responses and emission factors to Australian trees within MEGAN and ran the CSIRO Chemical Transport Model for three summertime campaigns in Australia. Compared to the default model, the new temperature responses resulted in less isoprene emission in the morning and more during hot afternoons, improving the statistical fit of modelled to observed ambient isoprene. Compared to current conditions, an additional 2 ppb of isoprene is predicted in 2050 causing hourly increases up to 21 ppb of ozone and 24-hourly increases of 0.4  $\mu\text{g m}^{-3}$  of aerosol in Sydney. ~~This forecasted increase in ozone is~~ [A 550 ppm CO<sub>2</sub> atmosphere in 2050 mitigates these peak Sydney ozone mixing ratios by 4 ppb. Nevertheless, these forecasted increases in ozone are up to](#) one fifth of the hourly Australian air quality limit and suggests anthropogenic NO<sub>x</sub> should be further reduced to maintain healthy air quality in future.

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## 1 Introduction

Biogenic volatile organic compounds (BVOCs) are emitted by vegetation in response to external stressors such as heat, light and herbivory (Sharkey and Monson, 2017). There are hundreds of individual BVOCs all exhibiting different emission behaviours (e.g. with or without a light dependence), but the largest global flux of a single BVOC is isoprene (2-methyl-1,3-butadiene;  $C_5H_8$ ), with an estimated 440 - 600 Tg C per year (Guenther et al., 2012). Isoprene reacts rapidly in the atmosphere, contributing to ozone ( $O_3$ ) and secondary organic aerosol (SOA) formation. For cities surrounded by forests, BVOC emissions can dominate airsheds contributing to peak summertime ozone (Utembe et al., 2018), and early morning ozone spikes (Millet et al., 2016) if not quenched by the hydroxyl radical (OH) on the previous day.

In addition to the different environmental reasons a plant will emit BVOCs, plants emit their own unique signature of BVOCs with varying strengths, even amongst plants in the same genus. Native to Australia, eucalypt trees are amongst the highest BVOC emitters of any plant species (Benjamin et al., 1996) (Evans et al., 1982; Benjamin et al., 1996; Kesselmeier and Staudt, 1999), emitting isoprene constitutively and storing monoterpenes within oil reservoirs in the leaves (Brophy et al., 1991). However, very few of the 800 species in the *Eucalyptus* genus (Department of Agriculture and Water Resources, 2018) have been studied for emissions. This is problematic as biogenic modellers tend to base simulations on a few measurements which represent a fraction of the potential diversity of species and emission rates. For example, the *Eucalyptus* isoprene emission factors for the Model of Emissions of Gases and Aerosols from Nature (MEGAN) were based on six studies, only one of which was conducted in Australia (see Emmerson et al. (2016)). The Australian study measured large differences of  $63 \mu g g^{-1} h^{-1}$  of isoprene between the lowest and highest emitting eucalypt species, with *E. globulus* showing the greatest emission rates (He et al., 2000). Natural occurrence of *E. globulus* is restricted to temperate south east Australia (including Tasmania).

Use of landscape emission factors (LEF) weighted by higher emitting trees have caused overpredictions in modelled isoprene (Emmerson et al., 2016, 2019a). As young leaves tend to emit more isoprene than older leaves, conducting emission measurements on saplings has been questioned (Street et al., 1997), although adult trees will contain a mixture of leaf ages. However, BVOC emission models such as MEGAN require isoprene emission rates to be determined at standard conditions of 303 K and  $1000 \mu mol m^{-2} s^{-1}$  photosynthetically active radiation (PAR) (Guenther et al., 2012). Measurements made at other temperatures and PAR fluxes need scaling to these standard conditions, which can introduce uncertainties of up to 20 % (He et al., 2000). The standard temperature and light level conditions are better provided for in a controlled greenhouse environment, which necessitates using saplings.

MEGAN describes the emission of BVOCs in terms of temperature, PAR, leaf area index, leaf age, soil moisture, and suppression via ambient  $CO_2$  concentrations. ~~The Whilst the~~ MEGAN parameterisations are ~~rightly based on fitted from~~ a wide range of ecosystem responses to environmental conditions, ~~there are spatial and temporal exceptions to these standards which are comprehensively reviewed by Niinemets et al. (2010)~~. Many studies have investigated impacts of climate change on isoprene by changing the inputs to MEGAN such as ambient temperatures and  $CO_2$  concentrations (e.g. Bauwens et al. (2018), and how land use might change the geographical extent of plant functional types (PFTs) (e.g. Arneth et al. (2011)), without changing the MEGAN parameterisations themselves. Here we report new controlled isoprene response measurements from

four eucalypt tree species, which show different temperature responses than assumed by MEGAN. We also use the controlled experimental conditions to impose a projected 2050 climate to investigate whether eucalypts growing in a warmer climate show a different temperature sensitivity of isoprene emissions than eucalypts growing in the current climate. Accounting for climate warming impacts on isoprene emission capacity provides a lens to study how air quality in Australia could be impacted in the future. Using a regional chemical transport model allows us to alter the dynamics of MEGAN to suit these new temperature responses for Australia.

This study aims to i) determine the temperature response of isoprene in four Eucalyptus species grown under two treatments representing current average summertime temperatures and a 2050 climate, and ii) use these measurements to determine the impacts of isoprene in a future climate on predicted levels of O<sub>3</sub> and SOA.

## 2 Methods

### 2.1 The MEGAN default temperature response

Guenther et al. (2012) defines the emission of BVOCs in terms of activity factors representing the environmental conditions described above. Here we are interested in studying the temperature response of isoprene,  $\gamma T$  (unitless):

$$\gamma T = E_{\text{opt}} \times \left[ \frac{C_{T2} \times \exp(C_{T1} \times x)}{(C_{T2} - C_{T1} \times (1 - \exp(C_{T2} \times x)))} \right] \quad (1)$$

where  $E_{\text{opt}}$  is the optimum emission point, and  $C_{T1}$  (95 kJ mol<sup>-1</sup>) and  $C_{T2}$  (230 kJ mol<sup>-1</sup>) are coefficients that fit the response to a range of ecosystems.

$$x = \left[ \frac{(1/T_{\text{opt}}) - (1/T)}{0.00831} \right] \quad (2)$$

where  $T$  is the temperature of the leaf (K) and 0.00831 is the gas constant in kJ K<sup>-1</sup> mol<sup>-1</sup>. The optimum temperature for emission in MEGAN,  $T_{\text{opt}}$  is calculated below.

$$T_{\text{opt}} = T_{\text{max}} + (0.6 \times (T_{240} - T_S)) \quad (3)$$

$$E_{\text{opt}} = C_{\text{eo}} \times \exp(0.05 \times (T_{24} - T_S)) \times \exp(0.05 \times (T_{240} - T_S)) \quad (4)$$

where  $T_{\text{max}}$  is 313 K,  $T_S$  is the standard leaf temperature (297 K), and  $T_{24}$  and  $T_{240}$  are the average leaf temperatures of the previous 24 and 240 hours, respectively.  $C_{\text{eo}}$  is an empirical coefficient of 2 for isoprene.

### 2.2 Experimental conditions

Four eucalypt species were chosen based on their prevalence in Australia, and in particular New South Wales (Table 1). ~~Two of the trees,~~ *E. camaldulensis* and *E. tereticornis* have a wide geographical representation within Australia, ~~having with~~ a latitudinal native growing range of 9-38 °S (Atlas of Living Australia, 2019) ~~-(supplementary figure S1).~~ *E. camaldulensis* is the most widely naturally distributed species of all eucalypts in Australia (Atlas of Living Australia, 2019). The native climatic

80 distribution range of ~~the other species~~, *E. botryoides* and *E. smithii* are restricted to the south east coastal regions. ~~We will use our new experimental data to revise the LEF maps for Australia, weighting the results according to the summed area of the four species (Table 1)~~All four species are forecast to exist in future, but only *E. camaldulensis* is predicted to expand its growing area by 2085 (González-Orozco et al., 2016).

Plant species can be classified as low (less than  $1 \mu\text{g g}^{-1} \text{h}^{-1}$ ), moderate ( $1\text{--}10 \mu\text{g g}^{-1} \text{h}^{-1}$ ) or high (greater than  $10 \mu\text{g g}^{-1} \text{h}^{-1}$ )  
85 isoprene emitters (Benjamin et al., 1996). Of the four eucalypts used in this study, *E. camaldulensis* and *E. tereticornis* are high isoprene emitters (Table 1), whilst *E. botryoides* is classed as moderate. The emission category of *E. smithii* is unknown. All tabulated measurements were scaled to the standard conditions from other temperatures and PAR.

Eighty trees (20 of each species) were grown from seed at Hawkesbury Institute for the Environment in Richmond, NSW. After eight weeks seedlings were transplanted into 6.9 L pots filled with alluvial soil and split randomly into two treatment  
90 groups, each containing 10 seedlings of each species. The first treatment group was grown for 85 days at an average daily temperature of 291 K (current climate) and the second treatment group was grown for 85 days at 294.5 K (future climate). In this time the seedlings put on vigorous growth and developed into  $\sim 1.5$  m tall saplings with plenty of leaves (see ~~photograph in supplementary supplementary figure S2~~). The future climate treatment represents temperature conditions in Australia in 2050 assuming the highest 8.5 Representative Concentration Pathway (RCP) the business as usual scenario where  $\text{CO}_2$  reaches  
95 940 ppm by 2100 (van Vuuren et al., 2011). The treatments maintained the diurnal variation of ambient temperature at 9 K. Further details on the growth conditions of these eucalypts are described in Aspinwall et al. (2019), prior to their study of how eucalypts respond to heatwave stress.

~~The measurements conducted on the future climate-grown trees provides the opportunity to study how isoprene emissions could change across Australia in a 2050 climate. This opportunity assumes the four eucalypt species exist in a 2050 climate and continue to have Australian coverage. Forward modelling suggests both *E. camaldulensis* and *E. tereticornis* will grow in Australia in 2085 (González-Orozco et al., 2016). We will use our new experimental data to revise the LEF maps for Australia, weighting the results according to the summed area of the four species (Table 1).~~

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### 2.3 Temperature response measurements

Leaf gas exchange measurements were made continuously with a LI-6400XT portable photosynthesis system (Li-Cor Inc.,  
105 Lincoln, NE, USA) connected to a Walz 3010-GWK1 leaf cuvette (maximum surface area for leaf  $140 \text{ cm}^2$ ; Heinz Walz GmbH, Effeltrich, Germany). The Walz cuvette was controlled via a PC using Walz software (GFS-Win v.3.47g).  $\text{CO}_2$  concentrations were set to 400 ppmv and the flow rate through the cuvette was set to  $700 \mu\text{mol s}^{-1}$ . Light was provided using Lumigrow Pro 325 LED growth lamps (LumiGrow, Novato, CA) positioned above the cuvette to provide  $1000 \mu\text{mol m}^{-2} \text{s}^{-1}$  PAR as measured by the LI-6400XT cuvette's light sensor. Leaf temperature was controlled using the Walz cuvette and was programmed to increase  
110 leaf temperature in 5 K steps from 293 K to 328 K in seven minute intervals to accommodate adjustment to new steady state values of photosynthesis at each temperature. This time corresponds to the duration of intermediate length sunflecks in plant canopies (Pearcy, 1990) and also results in a common, standardised heat dose for all the leaves (Niinemets and Sun, 2015). Basal emission rates are taken as the emission rate measured at  $1000 \mu\text{mol m}^{-2} \text{s}^{-1}$  PAR and 303 K.

After the gas exchange measurements, leaves were detached and their area measured using a LI-3100C leaf area meter (Li-Cor Inc.). Leaves were oven dried at 105 °C for 72 hours after which their dry weight was recorded.

Mixing ratios of isoprene by volume were determined using a high-resolution proton transfer reaction-mass spectrometer (PTR-MS, Ionicon GmbH, Innsbruck, Austria). The operating parameters of the PTR-MS were held constant during measurements, except for the secondary electron multiplier voltage, which was optimised before every calibration. The drift tube pressure, temperature and voltage were 2.2 hPa, 50 °C and 600 V. The parameter E/N was  $\sim 125$  Td ( $1.25 \times 10^{-15}$  V cm<sup>2</sup>) and the reaction time was  $\sim 100$   $\mu$ s. The count rate of H<sub>3</sub>O<sup>+</sup>·H<sub>2</sub>O ions was 1-2 % of the count rate of H<sub>3</sub>O<sup>+</sup> ions, which was  $5.0 - 5.5 \times 10^6$  s<sup>-1</sup>. Normalized sensitivities and isoprene volume mixing ratios were calculated through calibrations as described by Taipale et al. (2008) using 5 ppmv isoprene (Apel-Riemer Environmental Inc., Broomfield, CO) diluted in high purity nitrogen (BOC Ltd, Sydney, NSW). Protonated isoprene was detected by the PTR-MS as its molecular mass plus one (i.e. M + H+1 = 69). The duty cycle for each measurement period was 5 s.

Isoprene-temperature response measurements were replicated on five or six saplings of each species in each temperature treatment group ([supplementary figure S3](#)). The Solver program (Generalized Reduced Gradient nonlinear method, default settings; Microsoft Excel for Office 365; Microsoft Corporation, Redmond, WA, USA) was used to estimate four MEGAN coefficients, C<sub>T1</sub>, C<sub>T2</sub>, T<sub>max</sub> and C<sub>eo</sub> to minimise the difference between the result of Equation 1 and the measured temperature responses, for each tree species and growth temperature treatment. The basal emission rates for each species (in  $\mu$ g g<sup>-1</sup> h<sup>-1</sup>) were normalised to the average basal emission factor for that species and its growth temperature treatment. Normalising these data scales the actual emission rates and ensures they have a common basal emission factor of unity.

## 2.4 Observations of isoprene mixing ratios

Few measurements of ambient isoprene exist in Australia. Hourly observations made by Proton Transfer Reaction Mass Spectrometry are available for three summertime urban field campaigns near Sydney (Figure 1). These observations will be used to evaluate model predictions using our temperature response functions of isoprene emission. Isoprene observations are available from Bringelly in the January-February of 2007, SPS1 in Westmead in the February-March of 2011 (Keywood et al., 2019), and MUMBA in Wollongong in January-February of 2013 (Paton-Walsh et al., 2017, 2018). [Maximum \(and average\) measured temperatures were 308.9 K \(295.9 K\) for Bringelly, 310.0 K \(295.6 K\) for SPS1 and 317.2K \(295.3 K\) for MUMBA.](#) Climate projections for Australia forecast increases in average temperatures with an accompanying increase in the frequency of extreme heatwave days (Bureau of Meteorology and CSIRO, 2018). ~~There were several hours across these campaigns where maximum temperatures exceeded 303 K and 313 K. Temperatures above 303 K occurred for 37 hours at Bringelly, 29 hours in SPS1, and 27 hours during MUMBA. There were six hours during MUMBA where temperatures exceeded 313 K, with a maximum of 317.2 K on January 18th 2013.~~

## 2.5 The CSIRO Chemical Transport Model (C-CTM)

The C-CTM is a modelling framework designed to predict the atmospheric concentrations of gases and aerosols due to emissions, transport, chemical production and loss, and deposition. In addition to BVOCs, the framework has successfully predicted



pollen (Emmerson et al., 2019b), health effects from shipping (Broome et al., 2016) and air quality (Chambers et al., 2019). The C-CTM is driven by meteorology from the Conformal Cubic Atmospheric Model (CCAM, McGregor and Dix (2008)), taking boundary conditions from ERA-Interim. Four nested domains are used at spatial resolutions of 80 km, 27 km, 9 km and 3 km to downscale the atmospheric constituents over topography that increases with complexity at higher resolutions. The inner 3km domain contains 114 x 110 gridcells to encompass Sydney, Wollongong and the surrounding forested regions (Figure 1).

The model chemistry scheme is MOZART-T1 (Emmons et al., 2020) incorporating the latest research on isoprene oxidation pathways via additional radical production under low NO<sub>x</sub> conditions. The aerosol framework is a two-bin sectional scheme, processing organic species by the Volatility Basis Set (Shrivastava et al., 2008) and processing inorganic species via ISORROPIA\_II (Fountoukis and Nenes, 2007). The high and low NO<sub>x</sub> aerosol mass yields for the organic species, including isoprene, are provided by Tsimpidi et al. (2010).

Australia wide anthropogenic emissions come from an inventory based on human population density on a 10 km x 10 km grid resolution (updated from Physick et al. (2002)). Anthropogenic emissions for Sydney in the 3km domain are based on the most recent NSW inventory for the year 2008 (EPA NSW, 2012). The full canopy environment version of MEGAN2.1 (Guenther et al., 2012) was built into the C-CTM to calculate the biogenic emissions (Emmerson et al., 2016). Isoprene emissions,  $R$  in a given grid cell,  $xy$ , are predicted using LEF maps in combination with the land fraction,  $\chi$  occupied by 16 PFTs,  $j$ , using:

$$R = LEF_{x,y} \sum_{j=1}^{nPFT} (\gamma_{x,y} \times \chi_j) \quad (5)$$

Where  $\gamma$  represents the sum of all activity factors for light, temperature, soil moisture, leaf area index and leaf age. The  $\gamma$  for soil moisture is applied using data provided by the Soil-Litter-Iso model (SLI), as recommended by Emmerson et al. (2019a). Monthly leaf area index data come from MODIS MCD15A2 version 4.

A PFT map based on the ESA CCI Land Cover distribution for the year 2010 (ESA, 2016) was created. The ESA landcover data was used in conjunction with MODIS 44B (Vegetation Continuous Fields) product, level 5.1 for the year 2012 to provide the percentage tree, grass and shrub cover. Details on how these landcover data were aggregated or split into the 16 PFTs required by MEGAN2.1 are provided in the supplementary. Eucalypts fall under the broadleaf evergreen temperate tree category.

### 3 Results and discussion

#### 3.1 Temperature response results

The fitted temperature responses for each eucalypt tree species under both current and future climate growth conditions are stronger and shifted to higher leaf temperatures than the MEGAN2.1 default response (Figure 2). The peaks in current climate  $\gamma T$  are 40-90 % higher than default MEGAN, whilst the peaks in future climate  $\gamma T$  are 45-200 % higher. The position of the peaks are also shifted towards higher temperature optimums, by approximately 4-9 K. For the current climate growth

treatment results, running MEGAN with default settings would underestimate  $\gamma T$  and subsequently the isoprene emission at leaf temperatures greater than 303 K. MEGAN assumes that at growth temperatures lower than the standard conditions, the amplitude of the temperature response ( $E_{opt}$ ) is lowered and the peak of that response is shifted to a lower temperature ( $T_{opt}$ ). These new data show for all species studied, at each growth temperature, that this is not necessarily true. Our measurements also indicate that eucalypts have evolved to cope with the high Australian temperatures and can continue to protect against heat damage via isoprene emission until  $\sim 320$  K. Tree species with a wide geographical coverage such as *E. camaldulensis* may also be better adapted to surviving climate change (González-Orozco et al., 2016).

Each tree in each temperature treatment group produces a similar response (numbers of trees and their temperatures at maximum  $\gamma T$  given in Table 2). In the current climate-grown trees the temperature optimum in  $\gamma T$  is 317 - 318 K for *E. tereticornis* and *E. smithii* decreasing at higher leaf temperatures. Both *E. camaldulensis* and *E. botryoides* persist at high  $\gamma T$  until 328 K when measurements stopped. In the future climate-grown trees the  $\gamma T$  peak is also  $\sim 317$  K and there is a different response of *E. camaldulensis* and *E. botryoides* compared to the other species.  $\gamma T$  in *E. camaldulensis* increases steeply with increasing leaf temperature until 321.5 K thereafter decreasing sharply. This response is common amongst the five *E. camaldulensis* in the future climate treatment, although there is scatter around this fitted response. The *E. camaldulensis* result will dominate the weighted variables used in the modelling because of its larger geographic distribution (Table 1). We discuss the impact of this sharp downturn in  $\gamma T$  at high temperatures in section 3.3.

### 3.2 Isoprene emission rates

The basal isoprene emission rates (BER) in  $\mu\text{g g}^{-1} \text{h}^{-1}$  were measured at the standard 303 K and 1000  $\mu\text{mol m}^{-2} \text{s}^{-1}$  PAR (Table 2). As the current climate growth treatment represents current day climatic conditions, we only compare these with measurements made previously on the same species. Our *E. tereticornis* BER measurements are lower than those made by Nelson et al. (2000) and Jiang (2020); however our *E. camaldulensis* BER measurements are around 10  $\mu\text{g g}^{-1} \text{h}^{-1}$  higher than that listed by Benjamin et al. (1996), and our *E. botryoides* BER measurements are  $\sim 37$   $\mu\text{g g}^{-1} \text{h}^{-1}$  higher than that measured by He et al. (2000). He et al. (2000) used a mixture of young and mature leaves in their experiments. Our, which could be one explanation for the difference in emission rates as young leaves are expected to be higher emitters than older leaves in *Eucalyptus* (Street et al., 1997). However, as the growth conditions (particularly light and temperature) and measurement protocols between this study and He et al. (2000) were different (we directly measured BER with a leaf cuvette at 303 K and 1000  $\mu\text{mol m}^{-2} \text{s}^{-1}$  PAR while He et al. (2000) used a dynamic chamber and scaled emissions to 303 K and 1000  $\mu\text{mol m}^{-2} \text{s}^{-1}$  PAR using algorithms from Guenther et al. (1993)), it is difficult to undertake a direct comparison. However, our measurements put the four eucalypt species into the high emission category.

To create new isoprene emission factor maps suitable for the modelling, we convert the BERs into landscape emission factors ( $\text{LEF}_{\text{isop}}$ ). The average BER for each growth treatment is weighted according to their geographical areas in Table 1. BERs are then converted into LEFs using the leaf mass per unit area (LMA) in  $\text{g m}^{-2}$  and scaled with LAI in  $\text{m}^2 \text{m}^{-2}$ , similar to Emmerson

210 et al. (2018). The isoprene emission factor for trees in each temperature treatment is given by  $tree\_EF_{isop}$ :

$$tree\_EF_{isop} = BER \times LAI \times LMA \quad (6)$$

In the C-CTM, northern Australian vegetation is represented by broadleaf shrubs (30 to 40 %) and C4 grasses (50 to 80 % in some locations). If the isoprene emission factor maps are only based on the new eucalypt BERs, these are unlikely to be representative of shrubs and grasses. Here we ensure the non-tree fraction of grid cells in Australia are not impacted by these  
215 changes using the tree fraction ( $treefrac$ ) from the ESA product.

$$LEF_{isop} = (tree\_EF_{isop} \times treefrac) + (orig\_EF_{isop} \times (1 - treefrac)) \quad (7)$$

This leaves the fraction of original isoprene LEFs ( $orig\_EF_{isop}$ ) untouched for grass and shrub PFTs.

### 3.3 Impacts of changing $C_{T1}$ , $C_{T2}$ , $T_{max}$ and $C_{eo}$

Table 3 shows the results of fitting ~~the MEGAN variables according to the current and future climate growth treatment data.~~  
220 ~~These data are averages~~  $C_{T1}$ ,  $C_{T2}$ ,  $T_{max}$  and  $C_{eo}$  compared to the default MEGAN values. ~~These new fitted data are~~ for the four tree species in the experiment, weighted according to their coverage in Table 1. The new average LEFs from our four eucalypt species are ~~31-48~~ 31 - 48 % lower than the ~~average default average MEGAN~~ LEF we use in the base run for the 3km Sydney domain. ~~These reductions fit within the~~ Previous modelling showed that a 40 % estimated by reduction in isoprene was needed to better match the observations from our three field campaigns Emmerson et al. (2019a).

225 The value fitted for  $C_{T2}$  is very high (1158.36 kJ mol<sup>-1</sup>) in the future climate treatment compared with the ~~default and~~ current climate treatment ~~values~~ (167.11 kJ mol<sup>-1</sup>) and default MEGAN (230 kJ mol<sup>-1</sup>), due to the *E. camaldulensis* measurements in Figure 2. To assess whether  $C_{T2}$  should be re-fitted we examine the impacts of changing each of these variables one at a time using a MEGAN boxmodel designed in Jiang (2020). As the impacts of the new measurements are strongest at higher temperatures, we assume conditions from the hottest day in the MUMBA campaign (January 18th). The MEGAN boxmodel  
230 runs for 24 hours, and the results given as percentage changes to the maximum isoprene emission in Table 3. For the given fitted values on this day, the  $C_{T1}$  variable has the least and  $C_{eo}$  has the most impact on isoprene emissions. The high  $C_{T2}$  value in the future climate treatment ~~incurs a will not be refitted, as the incurred~~ 19 % decrease in isoprene, ~~which~~ is small compared with the 282 % increase caused by  $C_{eo}$ . Individually  $C_{eo}$  has the greatest impact on isoprene emissions but is regulated by increasing  $T_{max}$  when used in tandem with other variables. However, when all variables operate together the overall impact is  
235 an  $\sim 80$  % increase in isoprene emissions for both current and future climate growth conditions. Inclusion of the average LEF reduces the maximum isoprene emission by 7 % in the current climate treatment conditions and increases by 23 % in the future climate treatment conditions on the default.

### 3.4 Model experiment set-up

~~Six~~ Seven model experiments are defined (Table 4) and are run for the periods of the field campaigns described in section 2.4.  
240 We model the impacts of using the new current and future climate treatment temperature response variables separately from the

impacts of the new LEFs on atmospheric isoprene mixing ratios. For experiments 1 to 5, we use the same hourly meteorology, current day tree distribution maps and LAI datasets to drive the C-CTM. This allows us to separate the temperature effect in isoprene emissions from other influences which may change in a future climate. The intention is to investigate changes in isoprene emissions resulting from the temperature response results, not to combine these with future land-use changes and how the hourly meteorology will be impacted by climate change. However, in experiment 6 we use a simple delta-scaling approach to address how a future climate may impact the driving input temperatures to MEGAN.

We take the average change ( $\delta 2050$ ) in projected summertime surface temperatures for Australia under the RCP 8.5 scenario from eight models in the Coupled Model Intercomparison Project 5 (CMIP5) (for details see supplementary). Delta-scaling adds  $\sim 2$  K to the surface temperatures near Sydney. We only scale the surface temperature, thus experiment 6 is not a 2050 representation of the whole atmosphere. This restricts the use of the delta-scaled temperatures as a MEGAN input and not the temperature used for chemical reactions, as mass balance difficulties would occur by not also delta-scaling the pressure and air density through the height of the atmosphere. We estimate the reaction rate of isoprene with OH (calculated as  $2.54 \times 10^{-11} \exp(410/T)$  in MOZART-T1) would decrease by 1.7 % with the 3.5 K temperature rise between our current and future climate growth treatments.

~~Our implied future climate. The climate2050 run does not include the associated increases in CO<sub>2</sub> mixing ratios which would decrease isoprene emissions (Heald et al., 2009). When future changes in land use and CO<sub>2</sub> are considered, Sharkey and Monson (2014) suggest that net isoprene emissions will increase due to increasing temperature. Also, our future climate temperature treatment was to be consistent with our measurements that were also not conducted in a higher CO<sub>2</sub> atmosphere, so the model experiment is consistent. A 7th simulation assumes a 550 ppm CO<sub>2</sub> atmosphere on top of the delta-scaled surface temperatures, employing the Heald et al. (2009) method for calculating short and long term CO<sub>2</sub> activity factors,  $\gamma_C$ . Fixing the atmospheric CO<sub>2</sub> to 550 ppm reduces the isoprene emissions by 5 % in the short term and 13 % in the long term.~~

~~Delta-scaling adds  $\sim 2$  K to the surface temperatures near Sydney, resulting in an almost doubling of the number of hours above 303 K in the Bringelly and SPS1 field campaigns (74 and 53 hours, respectively), however none are pushed into the 313 K category. At MUMBA which already had six hours above 313 K, the number of hours in this category more than doubles to 13.~~

If the leaf temperature is varied within Equations 1-4 and  $\gamma T$  is multiplied by a normalised the LEF, the impacts of experiments 1-5 on isoprene emission start at about 283 K (Figure 3). Experiment 6 follows the FC- $\gamma T$ +LEF profile. Here, the new experimental current and future climate LEFs are normalised to by the default MEGAN LEF. The default MEGAN profile has a peak isoprene emission at 311 K. The CC- $\gamma T$  and FC- $\gamma T$  experiments cause the isoprene emission peak to shift to 324 K, with three times the default emission value. The sharp downturn in isoprene emission in the FC- $\gamma T$  and FC- $\gamma T$ +LEF experiments after 324 K are due to the high  $\gamma T$  of *E. camaldulensis* depicted in Figure 2. However, these results will not impact the C-CTM runs as no hourly temperature in our three field campaigns exceeds 317 K. Most of the impacts on the C-CTM runs will occur in the 288 - 308 K range. Whilst there is a very small decrease in the CC- $\gamma T$  response compared with the default MEGAN profile at temperatures less than 300 K, overall we expect more isoprene to be emitted in the CC- $\gamma T$  and FC- $\gamma T$  experiments over the default MEGAN profile. While it is intuitive from Figure 3 to expect less isoprene will be emitted in the CC- $\gamma T$ +LEF

and FC\_γT+LEF experiments over the base run (from Figure 3), this may not be the case due to spatial heterogeneity in the new current and future climate LEF maps. The LEFs used in Figure 3 are based on the domain spatial average value, however the LEFs in experiments 3 and 5 are based on the distribution of LAI from equation 6, whilst experiments 1, 2 and 4 use the original MEGAN LEF distribution. The ~~current and future climate average LEFs certainly~~ results from experiments 3 and 5  
280 certainly certainly show a sustained isoprene decrease below 314 K and 311 K respectively. Distance from source to receptor, transport and dilution will all impact results, and are determined by running the C-CTM.

### 3.5 C-CTM results

The C-CTM is compiled with changes to MEGAN implemented according to Table 3, ~~and~~ run for experiments 1-6 (Table 4). ~~The range in modelled and the isoprene time series is extracted at each field campaign site. The modelled mean~~ diurnal profiles of isoprene are then compared to the mean diurnal observations taken at each field campaign (Figure 4). Instrument calibrations/blanks are taken at least twice a day, incurring frequent regular gaps in observed isoprene. The CC\_γT variables only increase the isoprene mixing ratios when temperatures exceed 303 K (from Figure 3). This has changed the shape of the diurnal profiles of each field campaign in different ways, but generally the CC\_γT and CC\_γT+LEF experiments have increased ~~predicted statistical fits the diurnal modelled to observed r<sup>2</sup>~~ when compared with the ~~base run. r<sup>2</sup> between the~~  
285 base run and observations. The average modelled isoprene in the CC\_γT+LEF run is within ± 1 standard deviation of the observations 90 - 100 % of the time during Bringelly and SPS1, and 33 % during MUMBA which continues to exhibit high bias (supplementary figure S4). In MUMBA, the CC\_γT increases the isoprene mixing ratios above the base run between 11:00 and 17:00 AEDT in the heat of the day. Very hot temperatures during the day can often be accompanied by strong gusty winds from the Australian interior. The hottest campaign day, 18th January 2013 during MUMBA was associated with the high-  
295 est average hourly wind measurement of 8 m s<sup>-1</sup>. Hot and windy conditions would cause lots of sun-flecking within the tree canopy, causing very sudden temperature spikes on the leaf surface. Physiologically, the increased production of isoprene at these times may help mediate the impacts of these sudden during temperature and light spikes helps to maintain photosynthesis (Behnke et al., 2010) during times of mild stresses (Loreto and Fineschi, 2015), above and beyond leaf cooling via transpiration processes (Sharkey et al., 2008). High isoprene emitters can better survive prolonged heatwaves (Yanez-Serrano et al., 2019),  
300 although the Aspinwall et al. (2019) study on our four eucalypt species showed trees grown under future climate conditions suffered greater heatwave damage than the same species in current climate conditions.

During all campaigns the CC\_γT results have decreased the isoprene from the base runs in the morning between 08:00 and 11:00 AEDT, because these temperatures are less than 303 K where the γT are less than the default MEGAN profile (Figure 3). The CC\_γT+LEF experiments represent current day conditions, with roughly the correct magnitude (MUMBA excepted)  
305 of predicted isoprene and best statistical fit compared with the observations. The FC\_γT+LEF experiment has produced more daytime isoprene than the base run contrary to the prediction in Figure 3, because the distribution of isoprene LEFs near the field campaign sites is different to the default MEGAN LEFs. The climate2050 experiment adds between 110 ~~+170~~ - 170 % more isoprene during the day, or approximately 2 ppb. The addition of a higher CO<sub>2</sub> atmosphere has reduced the daytime isoprene by 15 - 26 % from the climate2050 run, across the three campaigns.

310 The MUMBA and SPS1 base diurnal profiles show too much isoprene in the model overnight compared to observed mean values, particularly in the period midnight to 06:00 AEDT. This is because there is more isoprene in the model atmosphere than was quenched by the OH radical before the OH production ceased at sundown. The isoprene becomes more concentrated at the surface because of the reduced boundary layer height; the apparent increase between midnight and 03:00 AEDT is not due to night-time isoprene emissions. [Conversely a slight rise in the model boundary layer at 04:00 AEDT in SPS1 causes dilution of](#)  
315 [the atmospheric isoprene.](#) While there are few measurements of isoprene during these pre-dawn periods, it is unlikely isoprene is present. Only when daytime isoprene is reduced in the CC\_γT+LEF experiment do we see the apparent night-time isoprene is decreased.

We investigate the spatial changes to isoprene, O<sub>3</sub> and biogenic SOA in an implied future by subtracting results from the CC\_γT+LEF experiment from the climate2050 experiment during the period of the SPS1 campaign (Figure 5). These emis-  
320 sions, mixing ratios and aerosol concentrations represent campaign averages from SPS1. We also show the smaller differences found between the FC\_γT+LEF and CC\_γT+LEF runs. The climate2050 experiment causes up to 5.2 mg m<sup>-2</sup> h<sup>-1</sup> in isoprene emissions to the immediate north of Sydney (Figure 5d), but there are also increases in the north of Australia (Figure 5c). The largest changes of 15.8 ppb in isoprene occur in sparsely inhabited northern Australia (Figure 5g), and in urbanised pockets to the south and east, where Sydney is located. The urbanisation becomes important when the increased isoprene reacts with NO<sub>x</sub>  
325 in the atmosphere causing a peak 9 ppb increase to O<sub>3</sub> near Sydney with the climate2050 differences (Figure 5l). However, the FC\_γT+LEF differences (Figure 5i) show a 0.5 ppb decrease in O<sub>3</sub> in northern Australia via quenching by the additional isoprene. Few inhabitants reside in northern Australia, meaning O<sub>3</sub> production via anthropogenic NO<sub>x</sub> is minimised. Soil NO<sub>x</sub> emissions are low in northern Australia as agricultural practices largely occur in the south east and south west of Australia. The O<sub>3</sub> deficit is still visible in the very north east of Australia in the climate2050 difference run (Figure 5k). The increase in  
330 biogenic SOA occurs mainly in the north of Australia where up to 0.21 μg m<sup>-3</sup> more aerosol is predicted by the climate2050 experiment than the CC\_γT+LEF experiment (Figure 5o).

The size fraction of most secondary organic aerosol fits within the PM<sub>2.5</sub> classification, defined as particulate matter with an aerodynamic diameter less than 2.5 μm. Australia sets National Environmental Protection Measures (NEPMs) for PM<sub>2.5</sub> and O<sub>3</sub> to ensure a healthy standard of air quality for the population. The NEPM for O<sub>3</sub> is 100 ppb as a 1-hour average, and  
335 25 μg m<sup>-3</sup> as a 24-hour average for PM<sub>2.5</sub>, with a goal of reducing the PM<sub>2.5</sub> limit to 20 μg m<sup>-3</sup> by 2025. We examine the increases brought about by climate induced isoprene in the two cities impacted by most of these changes, Sydney and Darwin, in Australia's north (Figure 6).

The air quality index (AQI = NEPM/pollutant concentration x 100) in Sydney and Darwin is classed as 'very good' (AQI <33) for both pollutants ([years 2009 - 2014](#)), with an improving trend for O<sub>3</sub> but a declining trend for PM<sub>2.5</sub> (Keywood et al.,  
340 2016). Darwin is a small city, and the biogenic component of O<sub>3</sub> changes are less than 2 ppb. However peak O<sub>3</sub> in Sydney increases by 10 - 15 ppb as an hourly average in the FC\_γT+LEF differences, but by [12 - 17 ppb in the climate2050 γC differences and by](#) as much as 15 - 21 ppb in the climate2050 differences (Figure 6a,b). These increases represent 10 - 21 % of the O<sub>3</sub> NEPM, and show that by doing nothing (e.g. tree type and coverage or air quality policies do not change) and allowing the temperatures to rise, large cities will likely encounter more NEPM exceedances. The solution is not to remove

345 native trees as they provide social amenity and have cultural significance for indigenous populations. Rather, their emissions must be accommodated via atmospheric NO<sub>x</sub> reductions. ~~However new~~ New urban developments should consider the BVOC emission potential of ~~prospective~~ trees before planting (Paton-Walsh et al., 2019), taking into account that non or low emitting trees may not withstand climate induced heatwaves (Penuelas and Munne-Bosch, 2005).

The SOA from isoprene is a small fraction of the PM<sub>2.5</sub> limit (shown here as 24-hour averages), though of the BVOC aerosol yields, isoprene is not expected to dominate. The aerosol yields from monoterpenes are 10-20 times higher than the isoprene yield and the monoterpene emission would increase in a warming climate (not investigated here). The climate2050 differences (and climate2050  $\gamma$ C) show days with an increase of ~~0.41~~ 0.42  $\mu\text{g m}^{-3}$  in Sydney and ~~0.13~~ 0.14  $\mu\text{g m}^{-3}$  in Darwin (2 % and 1 % of the PM<sub>2.5</sub> 2025 NEPM, respectively).

#### 4 Conclusions

355 We have measured the isoprene emission response to controlled increases in temperature from four eucalypt species, two of which have a large geographical growing extent in Australia. The trees were grown in temperatures representing the current climate summertime conditions in Australia and in temperatures representing the projected summertime conditions of +3.5 K warming under the business as usual RCP 8.5 scenario. Climate projections for Australia forecast increases in average temperatures with an accompanying increase in the frequency of extreme heatwave days (Bureau of Meteorology and CSIRO, 360 2018). This will likely increase in the number of days above 303 K.

The current condition experiments demonstrated a change in the isoprene emission response to temperature as compared with the default parameterisation in MEGAN. This is not a surprise, as MEGAN is built to represent a range in ecosystem responses, but may go some way to explain why difficulties have been encountered when modelling isoprene in Australia previously. Both the current and future climate growth treatment temperature responses shifted the peak in  $\gamma$ T by 4-9 K, signifying that these four 365 eucalypt species were observed to continue emitting isoprene until well past the default maximum temperature for emission at 313 K. This suggests the eucalypts used in this study have evolved to protect against higher temperatures as expected with climate change.

Higher basal emission rates were measured in three of the eucalypt species in our experiment than have been previously measured. However, the conversion of these average weighted emission rates to LEFs for use in the C-CTM, resulted in a 370 lower average LEF than are currently being used in the base run. This is due to low biomass measured on our leaves, and because the isoprene emission factors from regions described as shrubs or grasslands were not altered. The spatial distribution of the new LEFs were based on the LAI distribution, different to the default MEGAN isoprene LEF map.

The model results using the new current climate growth temperature responses improved the statistical fits of the diurnal profiles compared to the measurements in average isoprene across our three field campaign periods. The overall magnitude of 375 the modelled profile was also brought into better agreement with observations in combination with the new current climate growth LEFs. MEGANv2.1 essentially works using a series of variables dependant on vegetation type and biogenic compound



emission traits, and the results here suggest that the four MEGAN variables altered in our experiments could also become ecosystem or location specific.

~~Despite our measurements being~~ Our measurements were conducted on sapling trees which may exhibit higher isoprene emissions than adult trees ~~when emission rates are expressed on leaf mass basis but not on a leaf area basis (Street et al., 1997).~~ Street et al. (1997) explained this through younger leaves having a higher specific leaf area than older leaves because eucalypts exhibit heterophylly (the foliage leaves on the same plant are of two distinctly different types). The apparent difference in emission rates between young and old leaves could be a consequence of morphology rather than biochemistry, so we expect the trend between the current and future climate growth emissions to be similar amongst trees of all ages. Our model experiments simulating isoprene emissions in a 2050 climate examined the differences between these runs and the CC\_γT+LEF experiment. ~~Two-Three~~ future experiments were conducted, the first using current day meteorology, ~~and~~ the second using a delta-scaled surface temperature change to projected 2050 summertime temperatures, and the third using a 550 ppm atmospheric CO<sub>2</sub> on top of the delta scaled temperatures. The FC\_γT+LEF experiment showed increases in isoprene emissions in the north of Australia, as well as closer to Sydney. These increases led to O<sub>3</sub> rising 10 - 15 ppb close to Sydney as a result of the increased isoprene, whilst decreasing in sparsely populated northern Australia through quenching by the additional isoprene. The climate2050 experiment showed much larger increases in isoprene, O<sub>3</sub> and biogenic SOA across Australia, tempered slightly by the addition of increased atmospheric CO<sub>2</sub>. Delta-scaling the surface temperatures was the simplest way of conducting future climate experiments. Future work should investigate getting a downscaled version of the 2050 atmosphere from CCAM which would provide the hourly meteorology throughout the atmosphere that the C-CTM requires.

The future is expected to bring increased temperatures, CO<sub>2</sub> and land use changes. Sharkey and Monson (2014) evaluated the isoprene trade-off in each of these scenarios and concluded the temperature effects would dominate. O<sub>3</sub> is a secondary product of isoprene oxidation, and is currently maintained at healthy levels in Australia. In order to maintain these levels, air quality policy should investigate methods to reduce anthropogenic NO<sub>x</sub> emissions in city regions to accommodate these climate change induced increases in BVOC emissions. In addition, tree planting efforts in new urban developments should also consider the BVOC emission potential of prospective trees.

*Code availability.* TEXT

*Data availability.* The LAI data product was retrieved from MCD15A2 version 4 from the online Data Pool, courtesy of the NASA Land Processes Distributed Active Archive Center (LP DAAC), USGS/Earth Resources Observation and Science (EROS) Center, Sioux Falls, South Dakota, [https://lpdaac.usgs.gov/data\\_access/data\\_pool](https://lpdaac.usgs.gov/data_access/data_pool)

405 *Code and data availability.* TEXT



*Sample availability.* TEXT

*Author contributions.* KME and MP devised the modelling study and wrote the manuscript. KME conducted the modelling. MP, MJA, SP and MGT conducted the experimental work. MJA, SP and MGT edited the manuscript.

*Competing interests.* TEXT

410 *Disclaimer.* TEXT

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**Table 1.** Geographic range size of each *Eucalyptus* species in Australia and the isoprene emission rate by dry leaf weight basis

Tree	Common name	Area (km <sup>h</sup> ) <sup>a</sup>	% weight	Emission category	Average emission $\mu\text{g g}^{-1} \text{hr}^{-1}$
<i>E. camaldulensis</i>	River red gum	6 040 600	86.32	high	16.6 <sup>b</sup> <del>32.5<sup>c</sup></del> <del>28.0<sup>c</sup></del> <del>32.5<sup>d</sup></del>
<i>E. tereticornis</i>	Forest red gum	792 575	11.32	high	32.7 <sup>de</sup> 38.2 <sup>ef</sup>
<i>E. smithii</i>	Blackbutt peppermint	95 750	1.37	unknown	-
<i>E. botryoides</i>	Bangalay	74 175	1.06	moderate	5.3 <sup>b</sup>

<sup>a</sup> Species area in 2014 (from González-Orozco et al. (2016)). <sup>b</sup> He et al. (2000). <sup>c</sup> ~~Benjamin et al. (1996).~~ ~~Karkik and Winer (2001)~~ <sup>d</sup> ~~Benjamin et al. (1996).~~ <sup>de</sup> Nelson et al. (2000). <sup>ef</sup> Jiang (2020), from trees growing in ambient CO<sub>2</sub> concentrations.

**Table 2.** Average isoprene basal emission rates (BER), leaf mass per unit area (LMA) and temperature at maximum  $\gamma$ T from each pool of trees under the current and future climate growth conditions. Values in brackets are standard deviations. Data in right-hand column is derived from model fits.

Treatment	Species	No. of trees	BER, $\mu\text{g g}^{-1} \text{hr}^{-1}$	LMA, $\text{g m}^{-2}$	Temp at max $\gamma$ T, K
current climate	<i>E. tereticornis</i>	6	29.14 (13.91)	61.53 (5.42)	317.8
	<i>E. smithii</i>	6	41.21 (17.31)	54.93 (13.71)	317.8
	<i>E. botryoides</i>	6	42.46 (23.64)	72.51 (15.25)	318.4
	<i>E. camaldulensis</i>	6	42.87 (22.87)	72.79 (6.14)	322.1
future climate	<i>E. tereticornis</i>	6	41.57 (28.08)	64.05 (9.58)	317.3
	<i>E. botryoides</i>	5	55.18 (27.27)	77.96 (12.55)	317.5
	<i>E. smithii</i>	6	61.61 (20.01)	58.08 (5.10)	317.0
	<i>E. camaldulensis</i>	5	66.95 (22.44)	73.18 (4.64)	321.5

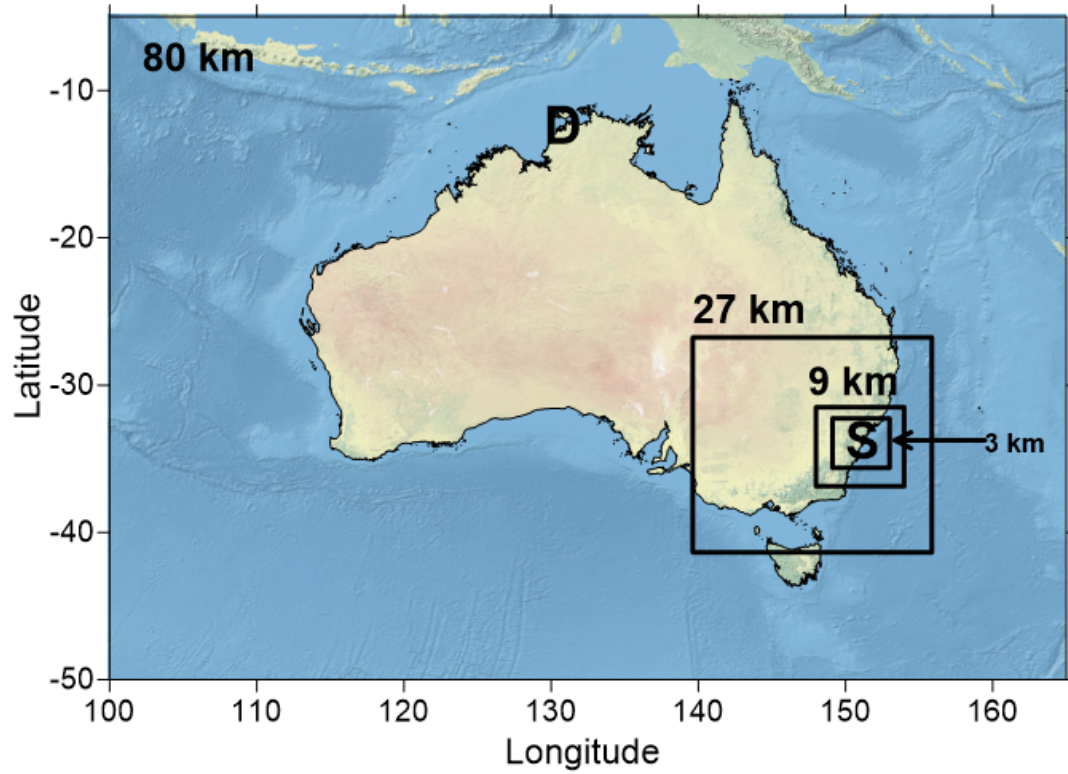
**Table 3.** Changes to MEGAN variables based on fitted data from current and future climate growth experiments. Percentages in brackets indicate change in maximum daily isoprene emissions due to change in variable. \*Value of average LEF from the inner 3 km domain.

	MEGAN2.1	current climate growth treatment	future climate growth treatment
Average LEF ( $\mu\text{g g}^{-1} \text{hr}^{-1}$ )	9491 *	4919 (-48%)	6585 (-31 %)
$C_{T1}$ ( $\text{kJ mol}^{-1}$ )	95	110.55 (-1 %)	75.04 (+1 %)
$C_{T2}$ ( $\text{kJ mol}^{-1}$ )	230	167.11 (+5 %)	1158.36 (-19 %)
$T_{\text{max}}$ (K)	313	325 (-55 %)	323 (-46 %)
$C_{\text{eo}}$	2	6.77 (+238 %)	7.69 (+282 %)
All variables without LEF		+81 %	+76 %
All variables + LEF		-7 %	+23 %

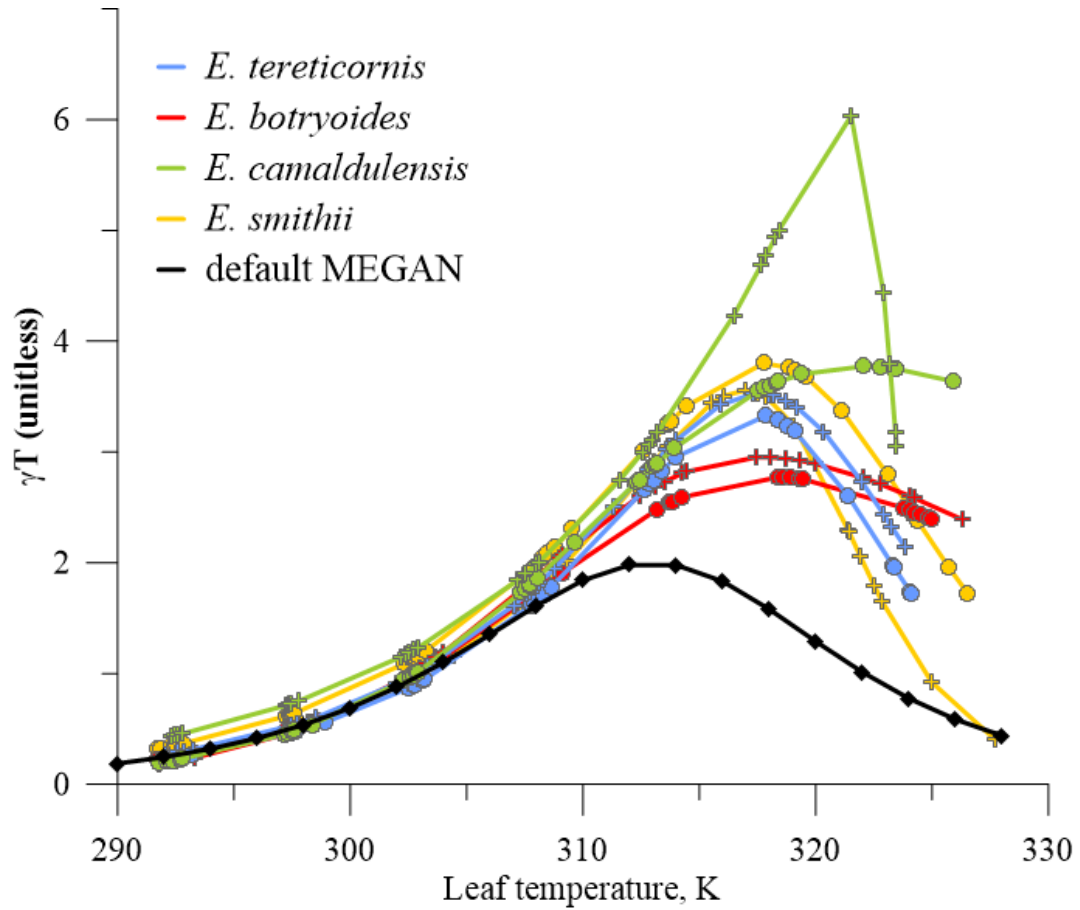


**Table 4.** Description of each model experiment. CC = current climate, FC = future climate.

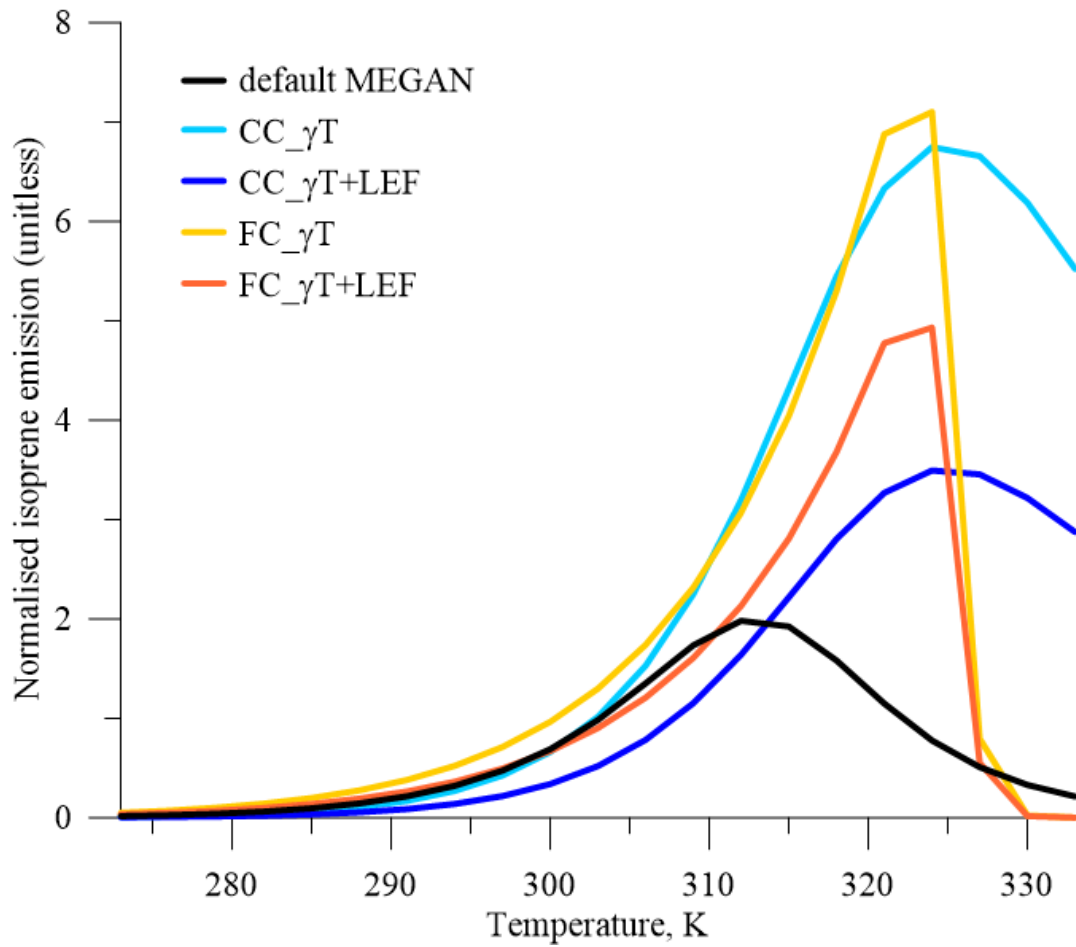
Experiment	Name	Emission factors	Temperature response	Meteorology used to drive MEGAN	$\gamma C$
1	Base	default	default	current	<u>x</u>
2	CC_ $\gamma$ T	default	fitted CC	current	<u>x</u>
3	CC_ $\gamma$ T+LEF	CC LEF	fitted CC	current	<u>x</u>
4	FC_ $\gamma$ T	default	fitted FC	current	<u>x</u>
5	FC_ $\gamma$ T+LEF	FC LEF	fitted FC	current	<u>x</u>
6	Climate2050	FC LEF	fitted FC	current + $\delta 2050$	<u>x</u>
<u>7</u>	<u>Climate2050_<math>\gamma C</math></u>	<u>FC LEF</u>	<u>fitted FC</u>	<u>current + <math>\delta 2050</math></u>	✓



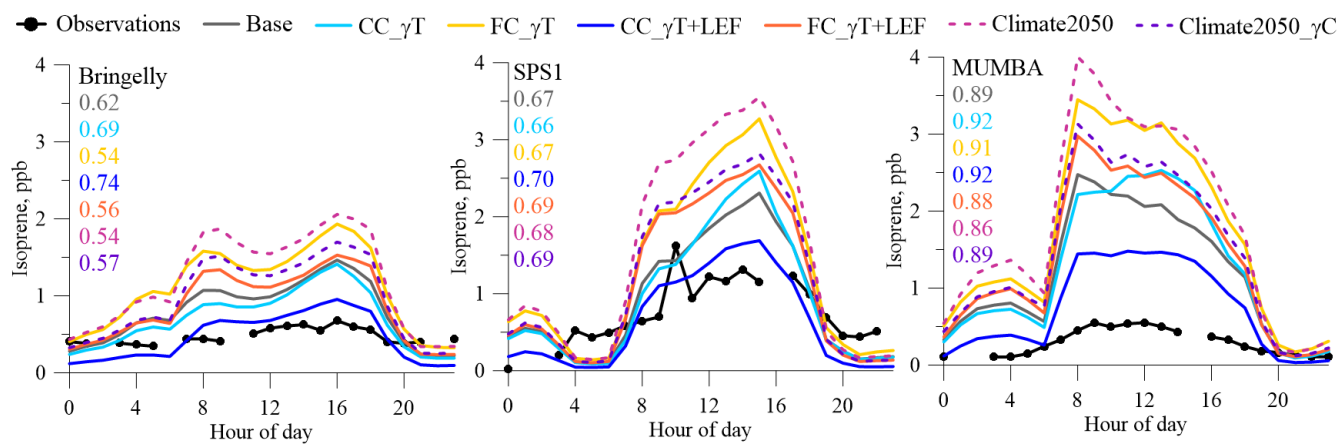
**Figure 1.** Map to show nests of model domains from 80 km Australia-wide to 3 km inner Sydney domain. S and D mark the locations of Sydney and Darwin, respectively.



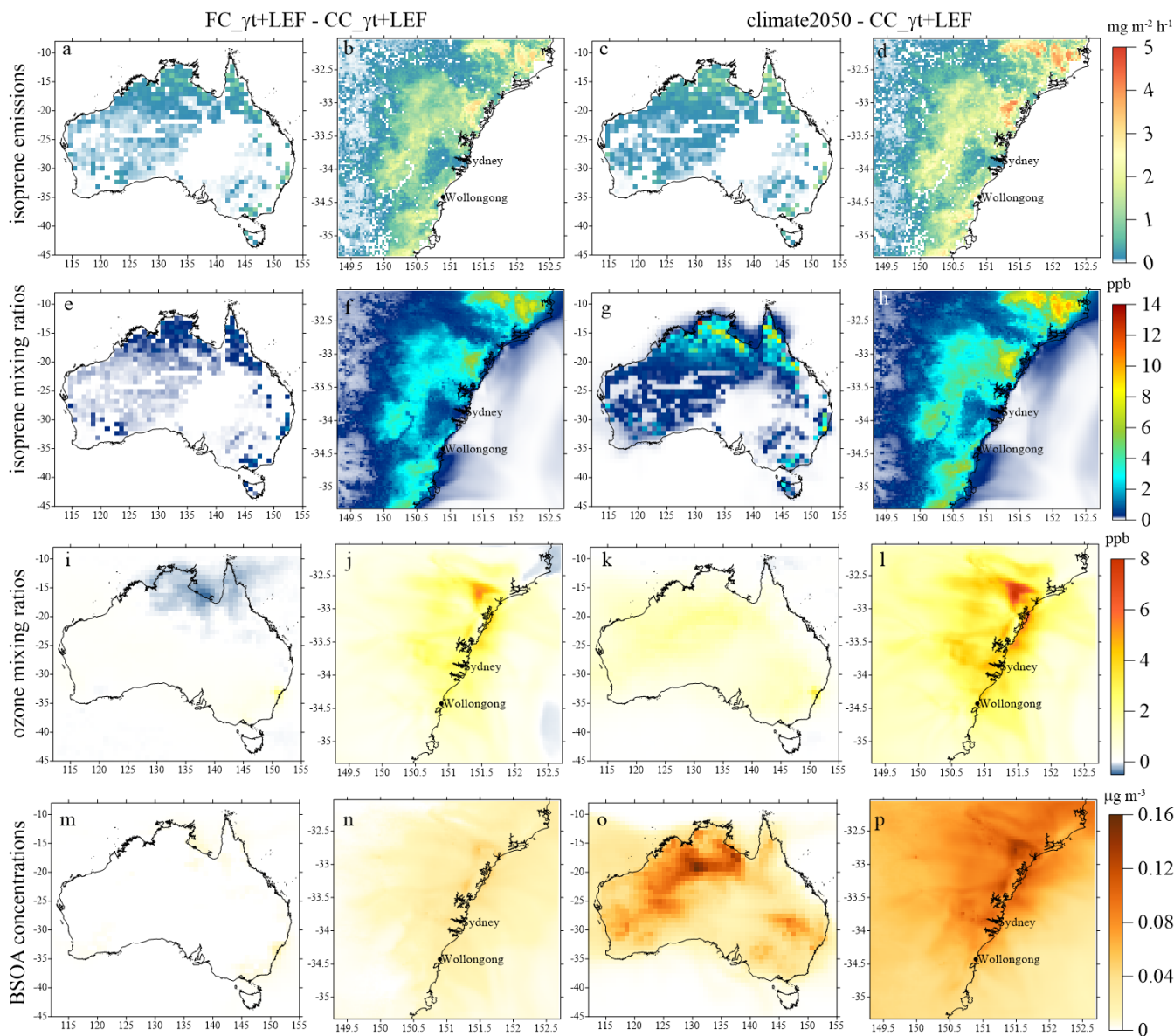
**Figure 2.** Comparison of  $\gamma T$  with leaf temperature calculated using default values in MEGAN to results from four eucalypt tree species under current climate (filled circles) and future climate (+ sign) growth conditions.



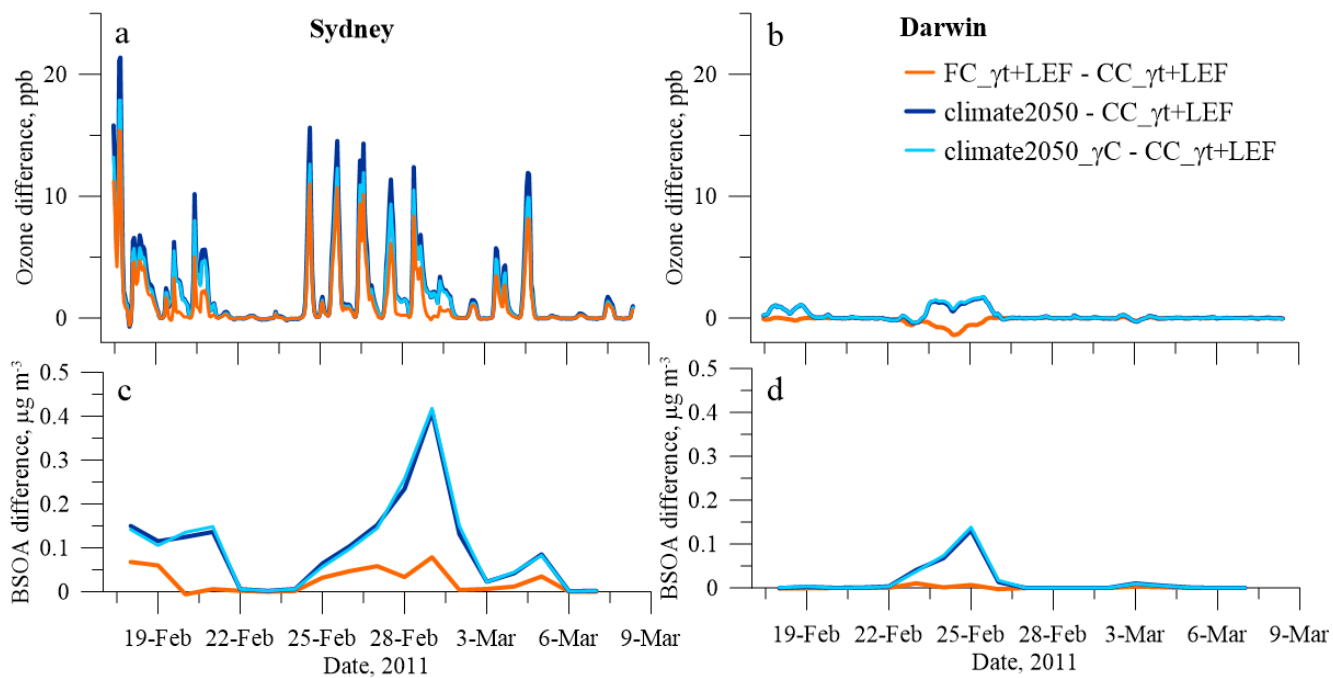
**Figure 3.** Impacts of new MEGAN variables on normalised isoprene emission rates at increasing ambient temperatures.



**Figure 4.** Average diurnal time series in isoprene mixing ratios incurred by the different model experiments at each field campaign site.  $r^2$  values between modelled and observed isoprene given in same colours as legend.



**Figure 5.** Difference between FC\_γT+LEF and CC\_γT+LEF runs (panels a, b, e, f, i, j, m and n) during the SPS1 campaign. The difference between the climate2050 runs and CC\_γT+LEF runs are shown in panels c, d, g, h, k, l, o and p. Left to right, panels a-d: Isoprene emission, panels e-h: isoprene mixing ratio, panels i-l: ozone mixing ratio and panels m-p: biogenic SOA concentration in Australia at 80 km and Sydney at 3 km domains.



**Figure 6.** Differences in hourly ozone (panels a and b) and biogenic secondary organic aerosol (panels c and d) [due to three 2050 experiments](#) at Sydney and Darwin during SPS1 [duration](#). [Left panels \(a and c\) show FC\\_γt+LEF - CC\\_γt+LEF](#) whilst [right panels \(b and d\) show climate2050 - CC\\_γt+LEF](#) [campaign period](#).