# Quantifying variability, source, and transport of CO in the urban areas over the Himalayas and Tibetan Plateau

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## 29 Abstract:

30 Atmospheric pollutants over the Himalayas and Tibetan Plateau (HTP) have potential 31 implications for accelerating the melting of glaciers, damaging air quality, water sources and 32 grasslands, and threatening climate on regional and global scales. Improved knowledge of the 33 variabilities, sources, drivers, and transport pathways of atmospheric pollutants over the HTP is 34 significant for regulatory and control purpose. In this study, we quantify the variability, source, and transport of CO in the urban areas over the HTP by using in situ measurement, GEOS-Chem model 35 36 tagged CO simulation, and the analysis of meteorological fields. Diurnal, seasonal, and interannual 37 variabilities of CO over the HTP are investigated with ~ 6 years (January 2015 to July 2020) of surface CO measurements in eight cities over the HTP. Annual mean of surface CO volume mixing 38 39 ratio (VMR) over the HTP varied over  $318.3 \pm 71.6$  to  $901.6 \pm 472.2$  ppby, and a large seasonal 40 cycle was observed with high levels of CO in the late autumn to spring and low levels of CO in 41 summer to early autumn. The diurnal cycle is characterized by a bimodal pattern with two 42 maximums in later morning and midnight, respectively. Surface CO VMR from 2015 - 2020 in

1 most cities over the HTP showed negative trends. The IASI satellite observations are for the first 2 time used to assess the performance of GEOS-Chem model for the specifics of the HTP. The GEOS-3 Chem simulations tend to underestimate the IASI observations but can capture the measured 4 seasonal cycle of CO total column over the HTP. Distinct dependencies of CO on a short lifetime 5 species of NO<sub>2</sub> in almost all cities over the HTP were observed, implying local emissions to be 6 predominant. By turning off the emission inventories within the HTP in GEOS-Chem tagged CO 7 simulation, the relative contribution of long range transport was evaluated. The results disclosed 8 that transport ratios of primary anthropogenic source, primary biomass burning (BB) source, and 9 secondary oxidation source to the surface CO VMR over the HTP varied over 35 to 61%, 5 to 21%, 10 and 30 to 56%, respectively. The anthropogenic contribution is dominated by the South Asia and East Asia (SEAS) region throughout the year (58% to 91%). The BB contribution is dominated by 11 12 the SEAS region in spring (25 to 80%) and the Africa (AF) region in July – February (30 - 70%). 13 This study concluded that main source of CO in urban areas over HTP is due to local and SEAS 14 anthropogenic and BB emissions, and oxidation sources, which differ from the black carbon that is 15 mainly attributed to BB source from Southeast Asia. The decreasing trends in surface CO VMR 16 since 2015 in most cities over the HTP are attributed to the reduction in local and transported CO 17 emissions in recent years.

#### 18 **1 Introduction**

19 The Himalayas and Tibetan Plateau (HTP), also named the 'Third Pole' (TP), is an important 20 region for climate change studies due to several reasons. Due to its unique feature for interactions among the atmosphere, biosphere, hydrosphere, and cryosphere, the HTP is referred to as an 21 22 important indicator of regional and global climate change (Pu et al., 2007; Yao et al., 2012; Zhang 23 et al., 2015). The HTP stores a large amount of ice masses on the planet and provides the headwater 24 of many Asian rivers which contribute water resource to over 1.4 billion people and it thus is referred 25 to as the 'Water Tower of Asia' (Xu et al., 2008; Immerzeel et al., 2010; Gao et al., 2019; Kang et 26 al., 2019). The glaciers and snowmelt over the HTP can potentially modify the regional hydrology, 27 contribute to global sea-level rise, and trigger natural hazards which may threaten the health and 28 wealth of many population (Singh and Bengtsson, 2004; Barnett et al., 2005; Immerzeel et al., 2010; 29 Kaser et al., 2010; Bolch et al., 2012; Yao et al., 2012; Gao et al., 2019). The HTP has an average 30 altitude of about 4000 m above sea level (a.s.l.), which highly elevates topography of the earth 31 system and imposes profound effects on global atmospheric circulation and climate change (Ye and 32 Wu, 1998; Wu et al., 2012; Zhang et al., 2015; Kang et al., 2019). The HTP is also of great interest 33 for model validation, since the extreme climate conditions and the variability between clean and 34 polluted conditions in the region are a challenge for current chemical transport models (Ye et al., 35 2012; Kopacz et al., 2011; Zhang et al., 2015).

36 Since the population level is very low, the HTP has long been regarded as atmospheric 37 background with negligible local anthropogenic emissions (Yao et al., 2012; Kang et al., 2019). 38 However, the HTP is surrounded by East Asia and South Asia which include many intensive 39 anthropogenic and natural emission source regions (Zhang et al., 2015; Kang et al., 2019). The 40 transport of polluted air masses from the highly populated area in northern India with its industry 41 and agriculture can have a strong impact on the composition of the atmosphere (Cong et al., 2009; 42 Kang et al., 2019). Furthermore, the Asian monsoon has a strong influence on the dynamics and 43 transport pathways in the HTP (Zhang et al., 2015; Kang et al., 2019). Reanalysis results based on

glacial ice cores and lake sediments have revealed distinguishable anthropogenic disturbances from 1 Asian emissions since the 1950s (Wang et al., 2008; Cong et al., 2013; Zhang et al., 2015; Kang et 2 3 al., 2016; Kang et al., 2019). Convective transport around the HTP areas has also been verified by 4 satellite observations, chemical transport model simulations, flask sampling analyses, and in situ 5 measurements of some key atmospheric constituents. These atmospheric constituents include 6 carbon monoxide (CO) (Park et al., 2007a; Park et al., 2007b), methane (CH<sub>4</sub>) (Xiong et al., 2009), 7 hydrogen cyanide (HCN) (Randel et al., 2010), polyacrylonitrile (PAN) (Zhang et al., 2009; 8 Ungermann et al., 2016; Xu et al., 2018), ozone  $(O_3)$  (Yin et al., 2017; Xu et al., 2018), and aerosol 9 (Cong et al., 2007; Cong et al., 2009; Cong et al., 2013; He et al., 2014; Zhang et al., 2015; Zhu et 10 al., 2019; Li et al., 2021; Gul et al., 2021; Thind et al., 2021). Furthermore, urbanization, 11 industrialization, land use, and infrastructure construction over the HTP have expanded rapidly in 12 recent years, which could also emit air pollutants into the atmosphere (Ran et al., 2014; Yin et al., 13 2019b).

14 The ecosystem over the HTP is sensitive and fragile under the extreme alpine conditions. These 15 exogenous and local atmospheric pollutants have potential implications for accelerating the melting 16 of glaciers, damaging air quality, water sources and grasslands, and threatening climate on regional 17 and global scales (Pu et al., 2007; Xu et al., 2009; Yao et al., 2012; Kang et al., 2016; Yin et al., 18 2019a; Yin et al., 2020; Yin et al., 2019b). Efforts have been made to understand the variabilities of 19 atmospheric pollutants over the HTP. However, due to the logistic difficulties and poor accessibility 20 of the vast HTP, most studies are based on episodic measurements in specific regions or at widely 21 dispersed sites (Kang et al., 2019). An inter-comparison of these data and deductions may show 22 large inconsistencies and uncertainties because the reported individual studies have often relied on 23 different instruments and techniques (Kang et al., 2019). In addition, most previous studies have 24 often concentrated on burdens, sources and transport of carbonaceous aerosols (including organic 25 carbon (OC) and black carbon (BC)) over the HTP, but the studies on gaseous pollutants are limited 26 (Cong et al., 2007; Cong et al., 2009; Cong et al., 2013; He et al., 2014; Zhang et al., 2015; Zhu et 27 al., 2019; Li et al., 2021; Gul et al., 2021; Thind et al., 2021). As a result, the variabilities, sources, 28 drivers, and transport pathways of atmospheric pollutants over the HTP are still not fully understood.

29 <u>CO</u> is one of the most critical atmospheric pollutant which not only threatens human health but 30 also plays a vital role in atmospheric chemistry (Zhang et al., 2019; Zheng et al., 2019). CO has a long atmospheric residence time of a few months and is therefore established as a key tracer for air 31 32 pollution and transport in the atmosphere (Holloway et al., 2000; Zheng et al., 2019). Natural 33 sources such as biomass burning (BB) and anthropogenic sources such as vehicle exhausts, 34 industrial activities, and coal combustions can emit CO directly into the atmosphere (Stremme et 35 al., 2013; Fisher et al., 2017). These CO emissions are mainly attributed to incomplete combustion 36 (Holloway et al., 2000; Stremme et al., 2013). Furthermore, the atmospheric oxidation of methane 37 (CH<sub>4</sub>) and numerous nonmethane VOCs (NMVOCs) provides additional important sources of 38 atmospheric CO (Fisher et al., 2017). The major CO sink in the troposphere is oxidation via reaction 39 with hydroxyl radicals (OH). Since CO is heavily involved in the relationship between atmospheric 40 chemistry and climate forcing, it is crucial to investigate its atmospheric burden, variability, and 41 potential driver over the HTP. CO over the HTP may originate from various source regions and 42 sectors, improved knowledge of their relative contributions to CO variability over the HTP is also significant for regulatory and control purpose. Furthermore, an investigation of CO pollution can 43 44 complement current atmospheric investigation over the HTP since the chemical characteristic,

1 climate forcing, and <u>deletion</u> of CO is different from the well-established carbonaceous aerosols.

2 In this study, we quantify the variability, source, and transport of CO in the urban areas over 3 the HTP by using in situ measurement, GEOS-Chem model tagged CO simulation, and atmospheric 4 circulation pattern techniques. Diurnal, seasonal, and interannual variability of CO over the HTP 5 are investigated with multiyear time series of surface CO measurements in eight cities over the HTP. 6 The performance of GEOS-Chem full-chemistry model for the specifics of the HTP is first assessed 7 with the concurrent satellite observations. The GEOS-Chem model is then run in a tagged CO mode 8 to quantify relative contribution of long range transport to the observed CO variability over the HTP. 9 The three-dimensional (3D) transport pathways of CO originated in various source regions and 10 sectors to the HTP are finally determined by the GEOS-Chem simulation, back trajectories analysis 11 and atmospheric circulation pattern. Only few studies have investigated the burden and variability 12 of CO over the HTP (Ran et al., 2014; Yin et al., 2019a). These studies uniformly focused on the 13 most developed regions in Lhasa, and did not analyze interannual trends and transport of CO. This 14 study not only expands the coverage of CO quantification over the HTP, but also provides insights 15 into the interannual trends, sources, and transport of CO in all urban areas over the HTP.

The next section describes site description, the surface in situ CO data and auxiliary data, the methodology used to estimate the interannual trend of surface CO, and the GEOS-Chem simulation used for source attribution. Section 3 reports the results for surface CO variability over the HTP on different time scales. Section 4 presents the results for GEOS-Chem model evaluation. Section 5 analyzes the results for source attribution using GEOS-Chem tagged CO simulation and the analysis of meteorological fields. We conclude the study in Section 6.

## 22 2 Methods and data

## 23 2.1 Site description

24 Surface in situ CO measurements in eight cities over the HTP are used in this study. The 25 locations of these cities are shown in Fig. 1 and summarized in Table 1. Ngari is located in the 26 western, Diqing and Qamdo is located in the eastern, and the rest cities are all in central eastern of 27 the HTP. Ngari, Shigatse, Lhasa, Shannan, and Nyinchi are adjacent to the Himalayas region, and 28 Naqu, Qamdo, and Diqing are relatively distant from the Himalayas region. Generally, these cities 29 represent the most developed and populated areas over the HTP. The altitude of these cities ranges 30 from 3.1 to 4.5 km a.s.l. and the population ranges from 110 to 770 thousand. The surface pressure 31 of these cities is about 600 hPa or less throughout the year (Table 1). Typically, all these cities are 32 formed at flat valleys with the surrounding mountains rising to more than 5.0 km a.s.l., and keep 33 continuous expansion and development over time. These cities are characterized by a typical climate 34 regime in high mountain regions, and is dry and cold in most of the year. Due to the high altitude 35 and thin air, incident solar radiations over these cities are stronger than those over other cities at the 36 same latitude around the globe (Ran et al., 2014).

General atmospheric circulation over these cities are typically influenced by three synoptic systems: the warm and wet air masses during the monsoon season in summer, the South Asian anticyclone that controls the upper troposphere and above, and the subtropical mid-latitude westerlies in winter (Yao et al., 2012; Ran et al., 2014; Yin et al., 2017). Inhibited by surrounding mountains, local mountain peak-valley wind systems facilitate the accumulation of atmospheric pollutants near the ground under low wind speed conditions (Kang et al., 2019).

#### 1 2.2 Surface CO data and auxiliary data

2 Routine in situ measurement of surface air qualities over the HTP started in 2015, which are 3 organized by the China National Environmental Monitoring Center (CNEMC) network funded by 4 the Chinese Ministry of Ecology and Environment (http://www.cnemc.cn/en/, last access: 22 March 5 2020). The CNEMC network has monitored six surface air pollutants (including CO, O<sub>3</sub>, NO<sub>2</sub>, SO<sub>2</sub>, 6 PM<sub>10</sub>, and PM<sub>2.5</sub>) at 23 sites in eight cities in Ngari, Lhasa, Naqu, Diqing, Shigatse, Shannan, 7 Nyingchi, and Oamdo over the HTP (Table 1). Each city has at least two measurement sites. Surface 8 CO volume mixing ratio (VMR) measurements at all sites are based on similar gas correlation filter 9 infrared analyzers (http://www.cnemc.cn/en/, last access: 22 March 2020). The hourly mean datasets 10 have covered the period from January 2015 to present for all measurement sites in the eight cities 11 (Table 1). We first applied filter criteria following that of (Lu et al., 2019) to remove unreliable 12 measurements. The resulting measurements at all measurement sites in each city are then averaged 13 to obtain a city representative dataset.

14 The 3D back trajectories calculated using HYbrid Single-Particle Lagrangian Integrated 15 Trajectory (HYSPLIT) model (http://ready.arl.noaa.gov/HYSPLIT.php, last accessed on 23 May 16 2020) are used to determine the transport trajectories (Wang, 2014; Draxler et al., 2020). The input 17 gridded meteorological fields were from the Global Data Assimilation System (GDAS-1) operated 18 by the US National Oceanic and Atmospheric Administration (NOAA) with a horizontal resolution 19 of 1° latitude × 1° longitude and 23 vertical grids from 1000 to 20 hPa (https://ready.arl.noaa.gov/ 20 gdas1.php, last accessed on 23 May 2020). We verified that the wind fields provided by GDAS-1 21 are in good agreement with those by the Goddard Earth Observing System-Forward Processing 22 (GEOS-FP) meteorological fields used in GEOS-Chem (Fig. S1). In this study, calculation and 23 analysis for all back trajectories are based on the TrajStat module (Wang, 2014; http://meteothink. 24 org/index.html, last accessed on 1 July 2020).

25 The monthly IASI/Metop-A CO dataset version 6.5.0 is used to evaluate the performance of 26 GEOS-Chem model for the specifics of the HTP. The IASI CO product is processed by EUMETSAT 27 Application Ground Segment using the Fast Optimal Retrievals on Layers for IASI (FORLI) 28 software (Hurtmans et al., 2012). The IASI CO retrievals are performed in the 2143-2181.25 cm<sup>-1</sup> 29 spectral range using the optimal estimation method and tabulated absorption cross sections at 30 various pressures and temperatures to speed up the radiative transfer calculation. A single a priori 31 profile is used in the retrieval scheme (Clerbaux et al., 2009). The temperature, pressure, humidity profiles and cloud fractions used in FORLI are those from the EUMETSAT Level 2 processor. Only 32 33 pixels associated with cloud fraction below 25 % are processed. The IASI CO product is a vertical 34 profile given as partial columns in moles per square meter in 18 layers between the surface and 18 35 km, with an extra layer from 18 km to the top of the atmosphere. The pressure levels associated 36 with retrieval layers are provided with the CO product. This IASI CO dataset also includes other 37 relevant information such as a general quality flag, the *a priori* profile, the total error profile, the air 38 partial column profile, and the averaging kernel (AK) matrix, on the same vertical grid, and the total 39 column and the associated total error. To balance the accuracy and the number of valid data over HTP, the IASI data within  $\pm 1^{\circ}$  latitude/longitude rectangular area around each city and with total 40 41 error of less than 15% are selected.

#### 42 2.3 Regression model for CO trend

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We have used a bootstrap resampling model to determine the seasonality and interannual

variability of surface CO VMR over the HTP. The resampling methodology follows that of
(Gardiner et al., 2008), where a 3<sup>rd</sup> Fourier series plus a linear function was used to fit multiyear
time series of surface CO VMR biweekly mean. All measurements are averaged by two weeks to
lower the residual and improve the fitting correlation. The regression model is expressed by Eqs. (1)
and (2):

6

$$Y^{meas}(t) = Y^{mod}(t) + \varepsilon(t)$$
(1)

7 
$$Y^{mod}(t) = A_0 + A_1 t + A_2 \cos\left(\frac{2\pi t}{365}\right) + A_3 \sin\left(\frac{2\pi t}{365}\right) + A_4 \cos\left(\frac{4\pi t}{365}\right) + A_5 \sin\left(\frac{4\pi t}{365}\right)$$
(2)

 $d\% = \frac{Y^{meas}(t) - Y^{mod}(t)}{Y^{mod}(t)} \times 100$ (3) where  $Y^{meas}(t)$  and  $Y^{mod}(t)$  represent the measured and fitted surface CO VMR time series,

9 where  $Y^{meas}(t)$  and  $Y^{mod}(t)$  represent the measured and fitted surface CO VMR time series, 10 respectively.  $A_0$  is the intercept,  $A_1$  is the annual growth rate, and  $A_1/A_0$  is the interannual trend 11 discussed below. In this study, we incorporated the errors arising from the autocorrelation in the 12 residuals into the uncertainties in the trends following the procedure of (Santer et al., 2008). The  $A_2$ 13  $-A_5$  parameters describe the seasonal cycle, *t* is the measurement time elapsed since January 2015, 14 and  $\varepsilon(t)$  represents the residual between the measurements and the fitted results. Fractional 15 differences of measured CO VMR time series relative to their seasonal mean values represented by 16  $Y^{mod}(t)$  were referred to as seasonal enhancements and were calculated as equation (3).

## 17 2.4 GEOS-Chem simulation

18 Two types of GEOS-Chem model simulations were involved in this study. GEOS-Chem model 19 version 12.2.1 (DOI:10.5281/zenodo.2580198) was first ran in a standard full-chemistry mode to 20 be evaluated by the IASI CO product. The GEOS-Chem model was then ran in a standard tagged 21 CO mode to quantify relative contribution of long range transport to the observed CO variability 22 over the HTP (Bey et al., 2001) (http://geos-chem.org, last access on 14 May 2020). The GEOS-23 Chem full-chemistry simulations were also used to provide OH fields and secondary CO production 24 rates from CH<sub>4</sub> and NMVOCs oxidation for subsequent GEOS-Chem tagged CO simulation. Both 25 types of simulations were driven by GEOS-FP meteorological fields with a downgraded horizontal 26 resolution of  $2^{\circ}$  latitude  $\times 2.5^{\circ}$  longitude and 72 vertical grids from surface to 0.01 hPa. Surface 27 meteorological variables and planetary boundary layer height (PBLH) were implemented in 1 hr 28 interval and other meteorological variables were in 3 hr interval. The time step used in the model 29 are 10 minutes for transport and 20 minutes for chemistry and emissions, as recommend for the 30 GEOS-Chem full-chemistry simulation at  $2 \times 2.5$  (Philip et al., 2016). The non-local scheme for the 31 boundary layer mixing process are described in Lin and McElroy (2010). The GEOS-Chem 32 simulation outputs 47 (tagged CO mode) or 72 (full-chemistry mode) vertical layers of CO VMR 33 concentration ranging from the surface to 0.01 hPa with a horizontal resolution of  $2^{\circ} \times 2.5^{\circ}$  and a 34 temporal resolution of 1 hr (Sun et al., 2020a). We spun up the model for one year (January 2014 to 35 January 2015) to remove the influence of the initial conditions. We only considered CO simulations 36 for the grid boxes containing the eight cities over the HTP.

Global fossil fuel and biofuel emissions were from the Community Emissions Data System inventory (Hoesly et al., 2018) which are replaced by regional emissions over the US by National Emission Inventory (NEI), Canada by Canadian Criteria Air Contaminant, Mexico by (Kuhns et al., 2005), Europe by European Monitoring and Evaluation Program (EMEP), East Asia and South Asia by MIX inventory (Li et al., 2017; Zheng et al., 2018; Lu et al., 2019), and Africa

by DICE-Africa inventory (Wiedinmyer et al., 2016). Global BB emissions are derived from Global 1 2 Fire Assimilation System (GFAS) v1.2 (Kaiser et al., 2012; Di Giuseppe et al., 2018). The soil  $NO_x$ 3 emissions were from (Hudman et al., 2010; Hudman et al., 2012). Biogenic emissions were from 4 the Model of Emissions of Gases and Aerosols from Nature (MEGAN version 2.1) inventory 5 (Guenther et al., 2012). Wet deposition followed that of (Liu et al., 2001) and dry deposition was calculated by the resistance-in-series algorithm (Wesely, 1989; Zhang et al., 2001). The photolysis 6 7 rates were obtained from the FAST-JX v7.0 photolysis scheme (Bian and Prather, 2002). A universal 8 tropospheric-stratospheric Chemistry (UCX) mechanism was implemented (Eastham et al., 2014). 9 In GEOS-Chem tagged CO simulation, the improved secondary CO production scheme of 10 (Fisher et al., 2017) was implemented, which adopts secondary CO production rates from CH<sub>4</sub> and 11 NMVOCs oxidation. The monthly mean OH fields and secondary CO production rates from CH<sub>4</sub> 12 and NMVOCs oxidation are archived from the full-chemistry simulation of this study. The GEOS-Chem tagged CO simulation includes the tracers of primary anthropogenic (fossil fuel + biofuel) 13 14 and BB sources, and secondary oxidations from CH<sub>4</sub> and NMVOCs. Descriptions of all these tracers

are summarized in Table 2 and the geographical definitions of all source regions are shown in Fig.
1.

## 17 **3 Surface CO variability over the HTP**

## 18 3.1 Diurnal cycle

Diurnal cycles of surface CO VMR over the HTP within the period of 2015 – 2020 are shown in Fig. 2. The surface CO magnitudes and the hour-to-hour variations in Naqu, Qamdo, and Diqing are higher than those in other cities in all seasons. Furthermore, the daily peak-to-trough contrast in Naqu, Qamdo, and Diqing are also larger than those in other cities. The highest surface CO hourly mean are typically observed in Naqu in all seasons except in the second half day (after 12:00 local time (LT)) in autumn and winter (September-October-November/December-January-February (SON/DJF)), when the highest surface CO values are observed in Qamdo.

26 Diurnal cycles of surface CO VMR in all cities generally show a bimodal pattern in all seasons. 27 For all cities, two diurnal maximums are generally observed during 9:00 to 11:00 LT in the daytime 28 and 21:00 to 23:00 LT in the nighttime in all seasons. The timings of the daytime diurnal maximum 29 in spring and summer (March-April-May/June-July-August (MAM/JJA)) in all cities are 1 to 2 30 hours earlier than those in SON/DJF (Table 3). But the timings of the nighttime diurnal maximum 31 in MAM/JJA in all cities are 1 to 2 hours later than those in SON/DJF. On average, the diurnal hour-32 to-hour variation of surface CO VMR over the HTP spanned a large range of -47.7% to 50.6% 33 depending on region, season, and measurement time. The diurnal patterns of CO in all cities over 34 the HTP were similar to those in other cities in China (Yin et al., 2019b; Zhao et al., 2016). Surface 35 CO VMR hourly mean in Naqu, Qamdo, and Diqing varied over  $455.8 \pm 257.8$  to  $1485.1 \pm 1104.7$ 36 ppbv, while other cities varied over  $256.4 \pm 177.1$  to  $650.0 \pm 430.7$  ppbv (Table 3). The Class 1 37 limit for the hourly mean CO concentration in China is 10 mg m<sup>-3</sup> (8732.1 ppbv) and all hourly 38 mean CO VMRs from 2015 - 2020 over the HTP were below this limit (http://www.cnemc.cn/en/, 39 last access: 22 March 2020).

## 40 3.2 Seasonal cycle

41 Seasonal cycle of surface CO VMR over the HTP within the period of 2015 to 2020 are shown
 42 in Fig. 3. As generally observed in most cities over the HTP, surface CO VMR showed clear seasonal

features: (1) high levels of surface CO VMR occur in the late autumn to spring and low levels of
surface CO occur in summer to early autumn; (2) the variations in the late autumn to spring are
larger than those in summer to early autumn; (3) seasonal cycles of surface CO VMR in most cities
show a bimodal pattern, i.e., a large seasonal peak occurs around November – December and a small
seasonal peak occurs around April – May.

6 Surface CO VMR monthly mean and month-to-month variations in Naqu, Qamdo, and Diqing 7 are higher than those in other cities in all seasons. Furthermore, the peak-to-trough contrast in Naqu, 8 Qamdo, and Diqing were also larger than those in other cities. Surface CO VMR monthly mean 9 over the HTP varied over a large range of  $206.8 \pm 93.5$  to  $1887.1 \pm 1132.0$  ppbv depending on 10 season and region (Table 3), where Naqu, Qamdo, and Diqing varied over 419.0  $\pm$  221.2 to 1887.1 11  $\pm$  1132.0 ppbv, and other cities varied over  $206.8 \pm 93.5$  to  $759.4 \pm 473.8$  ppbv (Table 3).

## 12 **3.3 Interannual variability**

13 Biweekly mean time series of surface CO VMR over the HTP from 2015 to 2020 along with the fitted results by using the regression model  $Y^{mod}(t)$  are shown in Fig. 4. Generally, the 14 measured and fitted surface CO VMR over the HTP are in good agreement with a correlation 15 16 coefficient (r) of 0.81 - 0.93. The measured features in terms of seasonality and interannual 17 variability can be reproduced by the regression model. Seasonal enhancements calculated as 18 equation (3) disclosed that large seasonal enhancements typically occur around November -19 December and April - May which correspond to the timings of the seasonal peaks for most cities. 20 The trend in surface CO VMR from 2015 to 2020 over the HTP spanned a large range of (-21.6  $\pm$ 21 4.5) % to  $(11.9 \pm 1.38)$  % per yr, indicating a regional representation of each dataset. Surface CO 22 VMR in Ngari, Lhasa, Shannan, Nagu, Oamdo, and Diging showed negative trends. The largest 23 decreasing trends were observed in Qamdo and Naqu, which showed decreasing trends of (-16.98 24  $\pm$  4.37) % and (-21.6  $\pm$  4.5) % per yr, respectively. Surface CO in Shigatse and Nyingchi showed 25 positive trends. A large increasing trend of  $(11.9 \pm 1.38)$  % per yr was observed in Shigatse.

Surface CO VMR annual mean over the HTP varied over  $318.3 \pm 71.6$  to  $901.6 \pm 472.2$  ppbv depending on year and region (Table 3), where Naqu, Qamdo, and Diqing varied over  $531.4 \pm 156.8$ to  $901.6 \pm 472.2$  ppbv, higher than those in other cities which varied over  $318.3 \pm 71.6$  to  $446.1 \pm 138.5$  ppbv (Table 3). The annual mean concentrations of surface CO over the HTP were compared with those from other cities in China. All cities over the HTP except Naqu and Qamdo can be ranked as a few of the top-level cities with the best air quality. Naqu and Qamdo were ranked as the middlelevel cities with fair to poor air quality (http://www.cnemc.cn/en/, last access: 22 March 2020).

33 4 Model evaluation over the HTP

34 The performance of the GEOS-Chem model has been evaluated with available observations 35 over various regions in China and surroundings in previous studies from different perspectives such 36 as surface O<sub>3</sub> concentration in urban regions over China (Lu et al., 2019), tropospheric CO column 37 over eastern China (Chen et al., 2009; Sun et al., 2020b) and Pacific (Yan et al., 2014), tropospheric 38 averaged HCHO concentration over eastern China (Sun et al., 2020a), stratospheric NO<sub>2</sub> partial 39 column (Yin et al., 2019a) and HCl partial column over eastern China (Yin et al., 2020). Generally, 40 GEOS-Chem is able to reproduce the absolute values as well as seasonal cycles of trace gases over 41 aforementioned regions. So far GEOS-Chem model evaluation over the complex topography and 42 meteorology of the HTP is not found in the literature. Here we first use IASI CO total column from

1 2015 to 2020 over the HTP to evaluate the model performance in the specifics of the HTP. As the 2 vertical resolution of GEOS-Chem is different from the IASI observation, a smoothing correction 3 was applied to the GEOS-Chem profiles. First, the GEOS-Chem CO profiles were downgraded to 4 the IASI altitude grid to ensure a common altitude grid. Since the IASI overpass time is at about 5 09:30 LT in the morning, only the GEOS-Chem simulations at 9:00 and 10:00 LT are considered. 6 The interpolated profiles were then smoothed by the monthly mean IASI averaging kernels and a 7 priori profiles (Rodgers, 2000; Rodgers and Connor, 2003). The GEOS-Chem CO total columns 8 were calculated subsequently from the smoothed profiles by using the corresponding regridded air 9 density profiles from the model. Finally, the GEOS-Chem total column time series were averaged 10 by month and compared with the IASI monthly mean data.

11 Correlation plots for the model-to-IASI data pairs in each region over the HTP are shown in 12 Fig. 5. Depending on regions, the GEOS-Chem simulations over the HTP tend to underestimate the 13 IASI observations by 9.2% to 20.0%. The largest GEOS-Chem vs. IASI differences occur in Qamdo 14 and Lhasa, with underestimations of 20.0% and 18.5%, respectively. The least GEOS-Chem vs. 15 IASI difference occurs in Nyingchi with an underestimation of 9.2%. These GEOS-Chem vs. IASI 16 differences over the HTP were mainly attributed to the underestimation of local emission inventories 17 and the coarse spatial resolution of the GEOS-Chem model grid cells. The amount of residential 18 energy use, including fossil fuels and biofuels used for cooking and heating, is not recorded for the 19 Tibet in current energy statistics yearbooks, therefore bottom-up inventories tend to underestimate 20 anthropogenic emissions over the HTP (Zheng et al., 2019). Meanwhile, emissions due to seasonal 21 crop residue burning over HTP are difficult to quantify accurately (Li et al., 2021). This can be 22 shown by the CO emission distribution over the HTP from the MEIC inventory in Fig. S2, which 23 shows that both the spatial distribution and seasonality of CO emission over the HTP are not in good 24 agreement with the in situ measurements. Besides, the coarse spatial resolution of the GEOS-Chem 25 simulations homogenizes CO concentrations within each  $2^{\circ} \times 2.5^{\circ}$  model grid cell. The simulation 26 results represent the homogenized concentrations in the grid box at the grid-mean elevation, which 27 could cause significantly bias near complex terrain (Yan et al., 2014). Especially, the studied regions 28 represent the most developed and populated areas over the HTP, which are surrounded by large areas 29 of rolling mountains with sparsely interspersed farms, pasture or residency. The horizontal transport 30 and vertical mixing schemes simulated by the GEOS-Chem model at coarse spatial resolutions are 31 difficult to match IASI observation with a ground pixel of a 12 km diameter footprint on the ground 32 at nadir (Clerbaux et al., 2009). Regional difference in CO levels could aggravate the inhomogeneity 33 within the selected GEOS-Chem model grid, and thus aggravate the difference between modeled 34 and measured CO concentrations. In addition, the difference between simulation and measurement 35 could be also associated with the uncertainties in meteorological fields, OH fields, and stratosphere-36 troposphere exchange (STE) scheme over HTP, which are known issues in the GEOS-Chem model 37 (Bey et al., 2001; Kopacz et al., 2011).

Though not perfect in reproducing the absolute values of the IASI observation, GEOS-Chem can capture the measured seasonal cycle of CO total column over the HTP with a correlation coefficient (*r*) of 0.64 to 0.82 depending on regions. In subsequent study, the GEOS-Chem model is used for investigating the influence of long range transport. We turn off all emission inventories within the HTP in the GEOS-Chem tagged CO simulation and assess the relative contribution of each source and geographical tracer. The relative contribution of each tracer is calculated as the ratio of the corresponding absolute contribution to the modelled total <u>concentration</u> amount. Taking this 1 ratio effectively minimizes the propagation of systematic model errors that are common to all tracers,

2 i.e., the uncertainties in meteorological fields, the vertical mixing and STE schemes, and the

3 mismatch in spatial resolution.

#### 4 **5** Source attribution

## 5 5.1 Local emission

6 The air quality in a city is influenced by local emission which is spatially differentiated by 7 energy consumption, economic development, industry structure, and population. All studied cities 8 over the HTP have achieved rapid economic and population growth in recent years (Ran et al., 2014; 9 Yin et al., 2019b). For example, Lhasa's gross domestic product (GDP) in 2018 was 29 times higher 10 than that of 2001, and the population had increased by more than 230 thousand in 17 years (Yin et 11 al., 2019b). In order to evaluate the influence of local emission, the relationship between in situ 12 measurements of NO<sub>2</sub> and CO is investigated. Correlation plots of surface CO versus NO<sub>2</sub> daily mean VMR time series provided by the CNEMC network from 2015 - 2020 in eight cities over the 13 14 HTP are shown in Fig. 6. The results show that NO<sub>2</sub> and CO concentrations were correlated in all 15 cities (r ranges from 0.49 - 0.86) throughout the year. The overall good correlations between these 16 two gas pollutants suggested common sources of  $\Delta NO_2$  and  $\Delta CO$  in these cities. As a short lifetime 17 species (a few hours), the emitted  $NO_2$  is heavily weighted toward the direct vicinity of local emission regions. As a result, local emissions are important sources of CO in all cities. However, 18 19 the slope  $\Delta NO_2/\Delta CO$  (ranges from 0.006 – 0.04) and the degree of the correlation in each city are 20 different, indicating energy consumption and CO emission rates in these cities are different, and 21 additional sources of CO could exist, e.g., from long range transport or oxidation from CH4 and 22 NMVOCs originating either nearby or in distant areas.

23 The emission from coal-burning for heating was thought to be the dominant sources of primary 24 gas pollutants in Lhasa in recent years (Ran et al., 2014; Yin et al., 2019b). A large portion of solely 25 source results in the highest correlation between NO<sub>2</sub> and CO concentrations in Lhasa. In contrast, 26 Qamdo, Naqu, and Diging are surrounded by alpine farmlands and pastures. Historically, post-27 harvest crop residue (e.g., highland barley straws and withered grass) was often burned by local 28 farmers to fertilize the soil for next planting season. As a fine fuel, post-harvest crop residue was 29 often burned directly in the field in large piles and smolder for weeks. These seasonal crop residue 30 burning behaviors typically occur in the cold season which could cause a high level of CO emission 31 in this period. Furthermore, local residents extensively use dry yak dung as fuel for cooking or 32 heating throughout the year which could elevate the background CO level in these regions. As a 33 result, these higher local sources might be an important factor explaining the higher CO magnitude 34 in these regions.

#### 35 **5.2 Long range transport**

Monthly mean contributions of anthropogenic, BB, and oxidation from long range transport to the surface CO VMR over the HTP are shown in Fig.7. All statistical results are based on GEOS-Chem tagged CO simulations by turning off the emission inventories within the HTP. Due to the influence of seasonally variable transport and magnitude of the regional emissions, the anthropogenic, BB and oxidation sources are all seasonal and <u>regionally</u> dependent. Generally, anthropogenic contributions in June – September and DJF are higher than those in the rest of the year. In contrast, high levels of oxidation contribution occur in JJA/SON and low levels of oxidation 1 contribution occur in MAM/DJF. For BB source, contributions in MAM/DJF are larger than those 2 in JJA/SON. Depending on season and region, relative contributions of anthropogenic, BB, and 3 oxidation transported to the surface CO VMR over the HTP varied over 35 to 61%, 5 to 21%, and 30 to 56%, respectively. The combination of anthropogenic and oxidation sources dominated the 5 contribution which varied over 80 to 95% with an average of 89% throughout the year.

6 After normalizing each regional anthropogenic contribution to the total anthropogenic 7 contribution, the normalized relative (NR) contribution of each anthropogenic region to the total 8 anthropogenic associated transport is obtained in Fig.8. The results show that the anthropogenic 9 associated transport is mainly attributed to the influence of anthropogenic sources in South Asia and 10 East Asia (SEAS). The NR anthropogenic contribution in SEAS ranges from 58% in DJF to 91% in 11 SON. In addition, moderate anthropogenic contributions from North America (NA) (10 to 27%), 12 Europe and Boreal Asia (EUBA) (4 to 12%), and rest of world (ROW) (4 to 10%) are also observed 13 in MAM/DJF. By using a similar normalized method, the NR contributions of each BB tracer and 14 oxidation tracer are obtained in Fig.9 and Fig. 10, respectively. The results show that large BB 15 contributions are from the Africa (AF) region in July – February (30 - 70%), the SEAS region in 16 MAM (25 to 80%), and the EUBA region in July – September (15 to 32%). Additional moderate 17 BB contributions are from the South America (SA) region in May - June and September - December 18 (9 to 14%), the Oceania (OCE) region in the second half of the year (5 to 15%), and the NA region 19 in September – December (8 to 19%). Depending on season and region, 45 to 67% of oxidation 20 contribution are attributed to CH<sub>4</sub> oxidation, and 32 to 55% of oxidation contribution are attributed 21 to NMVOCs oxidation. High-level NR contributions of CH<sub>4</sub> oxidation occur in the cold season 22 (November – March) and low-level NR contributions of CH<sub>4</sub> oxidation occur in the warm season 23 (April – October). The NR contributions of NMVOCs oxidation varied over an opposite mode to 24 that of CH<sub>4</sub> oxidation; they maximize in the warm season and minimize in the cold season. The 25 JJA/SON meteorological conditions that show stronger solar radiation, higher temperature, wetter 26 atmospheric condition, and lower pressure than those in DJF/MAM are more favorable for 27 increasing VOCs emissions from biogenic sources (BVOCs), which consolidates the fact that 28 contributions of NMVOCs oxidation in warm season are larger than those in cold season.

29 By minimizing the propagation of model errors that are common to all tracers (see section 4), 30 the major factors impacting the model interpretation are the uncertainties in emission inventories 31 and OH fields. The uncertainties in CO emission inventories mainly impact primary anthropogenic 32 and BB sources, and the uncertainties in CH<sub>4</sub> and VOCs emission inventories, and OH fields mainly 33 impact secondary oxidation sources. Additional factors that affect the generation and deplete 34 chemistry of CO or its precursors (e.g., uncertainties in emission inventories of other atmospheric 35 components, stratospheric intrusion of ozone and chemical mechanism, etc.) could also contribute 36 to the uncertainty of the interpretation. All these factors may be seasonal and regionally dependent. 37 A series of GEOS-Chem sensitivity studies might be able to quantify these uncertainties, but this is 38 beyond the scope of present work.

From section 5.1 and the model interpretation here, we can conclude that the main source of CO in urban areas over HTP is due to local and SEAS anthropogenic and biomass burning emissions, and oxidation sources. In contrast, black carbon in most of the HTP is largely attributed to Southeast Asian biomass burning, and locally sourced carbonaceous matter from fossil fuel and biomass combustion also substantially contribute to pollutants in urban cities and some remote regions, respectively (Cong et al., 2007; Cong et al., 2009; Cong et al., 2013; He et al., 2014; Zhang et al., 1 2015; Zhu et al., 2019; Li et al., 2021; Gul et al., 2021; Thind et al., 2021). Our study emphasized 2 the different origins of diverse atmospheric pollutants in the HTP.

#### **3 5.3 Transport pathways**

4 The 3D transport trajectories of CO originated in various source regions and sectors to the HTP 5 are identified as bellow. First, the GEOS-Chem tagged CO simulation is applied for determining seasonal NR contribution of each tracer (Figs. 8 and 9). For the tracer with a NR contribution of 6 7 larger than 30% at a specific time (hereafter enhancement time), the global CO distribution provided 8 by the GEOS-Chem simulation is applied to search for potential CO sources occurring before the 9 enhancement time within 15 days. Then, we generated a series of back trajectories with various 10 travel times to judge whether these CO emissions are capable of travelling to the measurement 11 region. For instance, with respect to each CO enhancement measured at a specific time, we 12 generated ten back trajectories arriving at 100 m above the ground but with different travel time 13 ranging from 3 to 15 days. If the back trajectories intersect a region where the GEOS-Chem 14 simulation indicates an intensive CO source and the travel duration is within  $\pm 2$  hr of the observed 15 enhancement, then this specific CO source could contribute to the observed enhancement over the 16 HTP. The transport trajectories for this CO source are finally determined. Meanwhile, GEOS-Chem 17 emission inventories are used to classify this CO source into anthropogenic or BB source. This CO 18 source is regarded as BB source if GEOS-Chem BB inventory indicates an intensive CO 19 enhancement. Otherwise, it is regarded as anthropogenic source.

20 Fig. 11 demonstrates travel trajectories of polluted air masses originated in AF, SEAS + OCE, 21 EUBA, and NA regions which arrived at Naqu (31.5°N) over the HTP through long range transport. 22 As the GEOS-Chem BB inventory shown, CO emissions from southern Africa during July – 23 September, central Africa during November - February, central Europe during July - November, 24 Siberia during June – September, and South Asia peninsula during March – May are dominated by 25 BB source. Other potential CO sources are dominated by anthropogenic emissions. Fig.12 shows 26 the spatial distribution of CO VMR along with the mean horizontal wind vectors at 500 hPa in 27 different seasons. Fig. 13 illustrates the latitude - height and longitude - height distributions of CO 28 VMR along with the 3D atmospheric circulation patterns in different seasons. The 3D transport 29 pathways of CO around the HTP are thus deduced as follows.

30 As indicated by the arrows in Fig. 12 and Fig.13, the strong surface cooling in DJF over the 31 HTP results in divergence and the formation of an enhanced local circulation cell, while in JJA air 32 masses converge toward the HTP from the surroundings triggered by the ascending of strongly 33 heated air masses over the HTP (Zhang et al., 2015). In DJF, the tropical easterlies are weak but the 34 mid-latitude westerlies extend to subtropics (~  $20^{\circ}$ N) near the surface and tropics (~ $10^{\circ}$ N) over 35 middle troposphere (Fig. 13). In the summer monsoon season, the atmospheric circulation patterns 36 around the HTP change dramatically and is dominated by the reversal of surface wind regime in the 37 tropics such as South China Sea, Bay of Bengal, and Arabian Sea (Fig.11 and Fig.13). Meanwhile, 38 the mid-latitude westerlies in JJA recede to the North Temperate Zone (north of 30°N) and the 39 westerly jet center shifts to about 40°N (from about 30°N in DJF). In JJA, the tropical region in the 40 south of the HTP is characterized by the strong easterlies in the upper troposphere and by the 41 southwesterly air flow in the lower troposphere (Fig. 13). The prevailing winds during the transition 42 seasons in MAM and SON are still westerlies (Fig. 13). These above seasonal atmospheric 43 circulation patterns control the CO transport pathway around the HTP. Nevertheless, the transported

1 CO scales to the HTP are also influenced by source location and strength, travel trajectory and 2 elapsed time (Yao et al., 2012; Zhang et al., 2015; Kang et al., 2019).

3 In the SON/DJF, a significant amount of CO from southern SEAS (anthropogenic source), 4 northern AF (BB source), western EUBA, and northern NA (anthropogenic source) can be 5 transported to the HTP along the westerlies in the dry winter monsoon conditions. CO originating 6 in distant regions such as western EUBA and NA reaches a high altitude (to 8 km) during the 7 transport (Fig. 11). However, CO from the densely populated and industrialized areas in eastern 8 China barely reaches the HTP because of strong removal along the transport pathways to the HTP which circles around the Northern Hemisphere along the westerlies during the winter monsoon 9 10 season (Fig. 12). In MAM, CO emissions from BB sources in SEAS region can be transported to 11 the HTP which is mainly triggered by deep convection followed by northward transport into the 12 mid-latitude westerlies (Liu et al., 2003)(Fig. 13). During the South Asian summer monsoon, the 13 local abundant wet precipitation can remove a large portion of SEAS originated CO but can still 14 affect southwest HTP (Fig. 13). Along strong southeasterly air flow in summer monsoon season, 15 CO from eastern China can be uplifted higher and transported more to the northeast HTP than that 16 in the DJF. In addition, large-scale atmospheric deep convection can loft CO from upwind source 17 regions (e.g., central SA, and Indonesia (within OCE region)) into higher altitudes, where it can be 18 transported to the HTP in SON or DJF. Generally, CO removals over the HTP in all seasons are 19 driven by atmospheric deep convection which lofts CO into higher altitudes or by westerlies which 20 transports local emissions far away (Fig. 13).

#### 21 5.4 Factors driving surface CO variability over the HTP

22 Temporal CO burden is dependent on the difference between the CO source and sink, which is 23 determined by the accumulated influence of local emission, transport, secondary generation, 24 environmental capacity, and OH oxidation capability. The environmental capacity is determined by 25 atmospheric self-clean capability, topography, deposition, and meteorological condition 26 (Hofzumahaus et al., 2009). Atmospheric self-clean capability refers to the capability of the 27 atmosphere in terms of depleting atmospheric pollutants through physical and chemical processes 28 (Rohrer et al., 2014). Generally, the vertical self-clean capability is positively correlated with the 29 PBLH and the horizontal self-clean capability is positively correlated with the wind speed (Rohrer 30 et al., 2014). The OH oxidation capability is positively correlated with temperature, radiation and 31 OH seasonality (Rohrer et al., 2014).

32 The bimodal pattern of diurnal cycles for surface CO VMR in urban areas over the HTP is 33 attributed to the following diurnal production and depletion processes. The thin atmosphere over 34 the HTP causes large temperature and radiation differences between day and night (Yin et al., 2017; 35 Kang et al., 2019). The PBLH and OH oxidation capability in the nighttime are much lower than 36 those in the daytime (Ran et al., 2014; Yin et al., 2017; Yin et al., 2019b). The CO emissions over 37 the HTP start to generate after sunrise and reach the daytime maximum during rush hours at 8:00 to 38 11:00 LT in the morning. The CO concentration is then decreasing as a result of depletion by reactions with OH to form O3 or transport far away (Ran et al., 2014; Yin et al., 2019b). Subsequently, 39 40 CO emissions start to generate again during rush hours at 16:00 to 19:00 LT in the afternoon, and 41 reach the nighttime maximum at 21:00 to 23:00 LT due to low PBLH and OH oxidation capability 42 in the nighttime (Ran et al., 2014; Yin et al., 2019b).

43

Similarly, the seasonal cycle of surface CO VMR in the urban areas over the HTP is determined

1 by the seasonal variability of CO source, environmental capacity, and OH oxidation capability. High 2 levels of surface CO VMR in the late autumn to spring can be attributed to low PBLH and OH 3 oxidation capability but high local and transported CO in the period, and vice versa for low levels 4 of surface CO VMR in summer to early autumn (Yin et al., 2019b). Specifically, local anthropogenic 5 CO sources (mainly heating activities) and crop residue burning behaviors in urban regions over the 6 HTP during the colder post-monsoon and winter months are higher than those in other seasons. 7 Meanwhile, the westerlies near the surface in SON/DJF are weaker than those in MAM/JJA, which 8 facilitate the accumulation of atmospheric pollutants (Fig. 13). Furthermore, high levels of CO are 9 observed in the late autumn to spring in neighboring SEAS countries due to intensive anthropogenic 10 emissions or BB practices (Kan et al., 2012; Tiwari et al., 2014; Liu et al., 2018; Gani et al., 2019). 11 These polluted air masses can transport to the HTP region and elevate the local CO level (Fig.13). 12 Thus, apart from local anthropogenic and BB emissions, these transported sources might be an 13 important factor explaining the high CO pollution in winter.

14 Since the crop residue burning emissions result in poor air quality that threatens local terrestrial 15 ecosystems and human health, Chinese government started to ban crop residue burning over China 16 since 2015 and henceforth the crop residue burning events over the HTP decreased dramatically 17 (Sun et al., 2020b) (http://www.chinalaw.gov.cn, last access on 19 June 2020). Meanwhile, the major 18 air pollutant emissions have decreased around the globe in recent years as a consequence of active 19 clean air policies for mitigating severe air pollution problems in the major anthropogenic emission 20 regions, such as China, India, Europe and America (Zheng et al., 2018; Sun et al., 2020a). 21 Furthermore, an elevated BB events in AF, SEAS, and OCE regions were observed at the beginning 22 of the studied years due to the El Niño Southern Oscillation (ENSO) in 2015 (Sun et al., 2020b). 23 All these factors probably drive a decreasing trend in surface CO VMR since 2015 in most cities 24 over the HTP. However, an overall increase in surface CO VMR in Shigatse and Nyingchi since 25 2015 indicated that the decrease in transported CO was overwhelmed by the increase in local CO 26 emissions as a result of the expansion of urbanization, industrialization, land use, and infrastructure 27 construction near the two cities.

#### 28 **6** Conclusions

29 In this study, we quantified the variability, source, and transport of CO in the urban areas over 30 the Himalayas and Tibetan Plateau (HTP) by using in situ measurement, GEOS-Chem model tagged 31 CO simulation, and the analysis of meteorological fields. Diurnal, seasonal, and interannual 32 variabilities of CO over the HTP are investigated with ~ 6 years (January 2015 to July 2020) of 33 surface CO measurements in eight cities over the HTP. Annual mean of surface CO volume mixing 34 ratio (VMR) over the HTP varied over  $318.3 \pm 71.6$  to  $901.6 \pm 472.2$  ppbv, and a large seasonal cycle 35 was observed with high levels of CO VMR in the late autumn to spring and low levels of VMR in 36 summer to early autumn. Surface CO VMR burdens and variations in Naqu, Qamdo, and Diqing 37 are higher than those in other cities in all seasons. The diurnal cycle is characterized by a bimodal 38 pattern with two maximums occurring around 9:00 to 11:00 local time (LT) in the daytime and 21:00 39 to 23:00 LT in the nighttime. The trend in surface CO VMR from 2015 to 2020 over the HTP 40 spanned a large range of (-21.6  $\pm$  4.5) % to (11.9  $\pm$  1.38) % per yr, indicating a regional 41 representatation of each dataset. However, surface CO VMR from 2015 to 2020 in most cities over 42 the HTP showed negative trends.

The IASI satellite observations are for the first time used to assess the performance of GEOS-

1 Chem full-chemistry model for the specifics of topography and meteorology over the HTP. 2 Depending on the region, the GEOS-Chem simulations over the HTP tend to underestimate the IASI 3 observations by 9.2% to 20.0%. Though not perfect in reproducing the absolute values of the IASI 4 observation, GEOS-Chem can capture the measured seasonal cycle of CO total column over the 5 HTP with a correlation coefficient (r) of 0.64 to 0.82 depending on regions. Distinct dependencies 6 of CO on a short life time species of NO2 in almost all cities over the HTP were observed, implying 7 local emissions to be predominant. By turning off the emission inventories within the HTP in GEOS-8 Chem tagged CO simulation, the relative contribution of long range transport was evaluated. The 9 results disclosed that transport ratios of primary anthropogenic source, primary biomass burning 10 (BB) source, and secondary oxidation source to the surface CO VMR over the HTP varied over 35 11 to 61%, 5 to 21%, and 30 to 56%, respectively. The anthropogenic contribution is dominated by the 12 South Asia and East Asia (SEAS) region throughout the year (58% to 91%). The BB contribution 13 is dominated by the SEAS region in spring (25 to 80%) and the Africa (AF) region in July – February 14 (30 – 70%). Additional important anthropogenic contributions from North America (NA) (10 to 15 27%) and Europe and Boreal Asia (EUBA) (4 to 12%) in spring and winter (MAM/DJF) are also 16 observed. Additional important BB contributions are from the EUBA region in July - September 17 (15 to 32%), the South America (SA) region in May – June and September – December (9 to 14%), 18 the Oceania (OCE) region in the second half of the year (5 to 15%), and the NA region in September 19 - December (8 to 19%). The decreasing trends in surface CO VMR since 2015 in most cities over 20 the HTP are attributed to the reduction in local and transported CO emissions in recent years.

21 This study concluded that the main source of CO in urban areas over HTP is due to local and 22 SEAS anthropogenic and BB emissions, and oxidation sources. In contrast, black carbon in most of 23 the HTP is largely attributed to Southeast Asian biomass burning, and locally sourced carbonaceous 24 matter from fossil fuel and biomass combustion also substantially contribute to pollutants in urban 25 cities and some remote regions, respectively. This study not only emphasized the different origins 26 of diverse atmospheric pollutants in the HTP, but also improved knowledge of the variabilities, 27 sources, drivers, and transport pathways of atmospheric pollutants over the HTP and provided 28 guidance for potential regulatory and control actions.

Data availability. Surface CO time series in all cities over the Himalayas and Tibetan Plateau (HTP) and GEOS-Chem tagged CO simulations in this study are available on request from Youwen Sun (ywsun@aiofm.ac.cn). The IASI product used in this study are available from https://iasi.aerisdata.fr/CO\_IASI\_C\_L3\_data/. The latest MEIC inventory is available from http://meicmodel.org/.

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 HY carried out the GEOS-Chem CO simulations. The rest authors contributed to this work via
 provide constructive comments.

36 *Competing interests.* The authors declare that they have no conflict of interest.

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#### Figures







Fig. 2. Diurnal cycles (local time (LT)) of surface CO VMR in four seasons over the Himalayas and Tibetan Plateau (HTP). Vertical error bar is  $1\sigma$  standard variation within that hour. Results are based on CO time series from 2015 to 2020 provided by the CNEMC network.



Fig. 3. Monthly mean of surface CO VMR over the Himalayas and Tibetan Plateau (HTP). Vertical error bar represent 1σ standard variation within that month. Results are based on CO VMR time series from 2015 to 2020 provided by the CNEMC network.



**Fig. 4**. Interannual variabilities of surface CO from 2015 to 2020 over the Himalayas and Tibetan Plateau (HTP). Green dots are biweekly mean of *in-situ* surface CO measurements. Vertical red error bar is  $1\sigma$  standard variation within the respective two weeks. The seasonality and interannual trend in each city fitted by using a bootstrap resampling model with a 3<sup>rd</sup> Fourier series (red dots) plus a linear function (black line) is also shown.



7Fig. 5. Correlation plots of monthly mean of CO total column over the Himalayas and Tibetan Plateau (HTP) for8GEOS-Chem model simulation against IASI observation. The average for GEOS-Chem simulation is performed at99:00 and 10:00 LT. The IASI dataset is selected within  $\pm 1^{\circ}$  latitude/longitude rectangular area around each city. The10blue lines are linear fitted curves of respective scatter points. The black dotted lines denote one-to-one lines.



1 2 3 4 5

Fig. 6. Correlation plots of surface CO versus NO2 VMR daily mean over the Himalayas and Tibetan Plateau (HTP). The black line is a linear least-squares fit of respective data. The linear equation of the fit and the resulting correlation coefficient (r) are shown. Both CO and NO<sub>2</sub> time series are provided by the CNEMC network. Vertical error bar is  $1\sigma$  standard variation within that day.



6

7 8 9 Fig. 7. Monthly mean contributions of anthropogenic, biomass burning (BB) and oxidation transport to CO over the Himalayas and Tibetan Plateau (HTP). Vertical error bar represent 10 standard variation within that month. See

Table 2 for description of each tracer.



**Fig. 8**. Monthly mean contributions of each geographical anthropogenic tracer to the total anthropogenic associated CO transport to the Himalayas and Tibetan Plateau (HTP). Vertical error bar is  $1\sigma$  standard variation within that month. See Table 2 for description of each tracer.



Fig. 9. Monthly mean contributions of each geographical BB tracer to the total BB associated CO transport to the
 Himalayas and Tibetan Plateau (HTP). Vertical error bar represent 1σ standard variation within that month. See
 Table 2 for description of each tracer.



**Fig. 10**. Monthly mean contributions of CH<sub>4</sub> and NMVOC oxidations to the total oxidation associated CO transport to the Himalayas and Tibetan Plateau (HTP). Vertical error bar represent  $1\sigma$  standard variation within that month. See Table 2 for description of each tracer.



Fig. 11. Travel trajectories of polluted air masses originated in AF (SON/DJF), SEAS & OCE (MAM/JJA), EUBA (SON/DJF), and NA (SON/DJF) that reached Naqu (31.5°N) through long range transport. For clarity, only few trajectories are selected for demonstration. Travel times are 13, 7, 10, and 14 days, respectively. Global surface CO distribution based on GEOS-Chem simulations are shown for 13, 7, 10, and 14 days prior to the arrival time, respectively. The deeper the color, the higher the CO concentration.



2 3 4 Fig. 12. Spatial distribution of CO VMR in the GEOS-Chem tagged CO simulations in different seasons, and the arrows represent the mean horizontal wind vectors at 500 hPa. The HTP and the studied regions are marked with a

blue outline and yellow dots, respectively. Meteorological fields are from the GDAS-1 data.



6 7 8 9 Fig. 13. The first row shows the latitude-height distributions of CO VMR averaged over 80-100° E in the tagged CO simulations in different seasons (corresponding to different columns). The white contours at intervals of 6 m s<sup>-1</sup> represent the westerly (solid) and easterly (dashed) mean meridional winds; the white area represents topography and the arrows represent the wind vectors (vertical velocity in units of  $10^{-4}$  hPa s<sup>-1</sup> and zonal wind in m s<sup>-1</sup>); the 10 studied regions are marked with yellow dots. The second row is similar to the first row except that the quantities are 11 on the longitude-height perspective averaged over 28-33° N. Here the white contours represent the southerly (solid) 12 13 and northerly (dashed) mean zonal winds, and the horizontal component of the wind vectors is meridional wind (m s<sup>-1</sup>). Meteorological fields are from the GDAS-1 data.

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#### 1 Tables

Table 1. Geolocations of measurement sites in eight cities over the HTP region. All sites are organised as a function

2 3 4 of increasing longitude. Population statistics are prescribed from the 2018 demographic data provided by National Bureau of Statistics of China.

Name	Longitude	Latitude	Altitude	Population	Number	Time period
	mean	mean	mean (km)		of sites	
Ngari	80.1°E	32.5°N	4.5	110,000	2	Jan. 2015 - present
Shigatse	88.9°E	29.3°N	4.0	770,000	3	Jan. 2015 - present
Lhasa	91.1°E	29.7°N	3.65	690,000	6	Jan. 2015 - present
Shannan	91.8°E	29.2°N	3.7	370,000	2	Jan. 2015 - present
Naqu	92.1°E	31.5°N	4.45	500,000	3	Jan. 2015 - present
Nyingchi	94.4°E	29.6°N	3.1	230,000	2	Jan. 2015 - present
Qamdo	97.2°E	31.1°N	3.26	700,000	3	Jan. 2015 - present
Diqing	99.7°E	27.8°N	3.38	410,000	2	Jan. 2015 - present

5 6 Table 2. Descriptions of all tracers implemented in the standard GEOS-Chem tagged CO simulation and the geographical definitions of all source regions.

Туре	Tracer	Description	Region	Location
Anthropogenic	anthNA	Fossil fuel + Biofuel CO	172.5°W - 50.0°W;	NA
		emitted over the North	24.0°N - 88.0°N	
		America		
	anthEUBA	Fossil fuel + Biofuel CO	17.5°W - 72.5°E; 36.0°N -	EUBA
		emitted over Europe and	45.0°N and 17.5°W -	
		Boreal Asia	172.5°E; 45.0°N - 88.0°N	
	anthSEAS	Fossil fuel + Biofuel CO	70.0°E - 152.0°E; 8.0°N -	SEAS
		emitted over South Asia	45.0°N	
		and East Asia		
	anthROW	Fossil fuel + Biofuel CO	112.5°W - 32.5°W; 56°S -	SA+AF+O
		emitted South America,	24°N and 17.5°W -70.0°E;	CE
		Africa and Oceania	48.0°S - 36.0°N and 70.0°E -	
			170.0°E; 90.0°S - 8.0°N	
Biomass	bbSA	Biomass burning CO	112.5°W - 32.5°W; 56°S -	SA
burning		emitted over South	24°N	
		America		
	bbAF	Biomass burning CO	17.5°W -70.0°E; 48.0°S -	AF
		emitted over Africa	36.0°N	
	bbSEAS	Biomass burning CO	70.0°E - 152.5°E; 8.0°N -	SEAS
		emitted over South Asia	45.0°N	
		and East Asia		
	bbOC	Biomass burning CO	70.0°E - 170.0°E; 90.0°S -	OC
		emitted over Indonesia	8.0°N	
		and Oceania		
	bbEUBA	Biomass burning CO	17.5°W - 72.5°E; 36.0°N -	EUBA
		emitted over Europe and	45.0°N and 17.5°W -	
		Boreal Asia	172.5°E; 45.0°N - 88.0°N	
	bbNA	Biomass burning CO	173°W - 50°W; 24.0°N -	NA
		emitted over North	88.0°N	
		America		
Oxidation	oxCH	CO chemically produced	global	global
		from CH4 oxidation		
	oxNMVOC	CO chemically produced	global	global
		from NMVOCs oxidation		

Name	Diurnal cycle		Seasonal cycle		Interannual variability	
	Hourly min.	Hourly max.	Monthly min.	Monthly max.	Annual mean	Trend
	(ppbv)	(ppbv)	(ppbv)	(ppbv)	(ppbv)	(% per yr)
Ngari	$319.2 \pm 150.6$	556.1 ± 311.0	$206.8 \pm 93.5$	$632.0 \pm 268.2$	$395.9 \pm 138.0$	-6.1 ± 1.39
_	(16:00)	(11:00)	(Aug.)	(Jan.)		
Shigatse	$256.4 \pm 177.1$	$464.9 \pm 253.1$	$274.4 \pm 98.2$	496.8 ± 365.3	$408.4 \pm 165.8$	$11.9 \pm 1.38$
_	(18:00)	(10:00)	(Sep.)	(Jan.)		
Lhasa	$334.5 \pm 101.0$	$650.0 \pm 430.7$	$375.8 \pm 128.6$	$759.4 \pm 473.8$	$446.1 \pm 138.5$	$-1.49 \pm 1.32$
	(17:00)	(22:00)	(Jul.)	(Dec.)		
Shannan	393.1 ± 230.3	$572.6 \pm 265.6$	$310.5 \pm 120.5$	665.8±280.2	409.8±147.0	$-3.68 \pm 1.42$
	(17:00)	(10:00)	(Oct.)	(Feb.)		
Naqu	757.3 ±411.6	$1485.1 \pm 1104.7$	607.1 ±284.3	$1539.4 \pm 1075.8$	$901.6 \pm 472.2$	$-21.6 \pm 4.5$
_	(17:00)	(23:00)	(Jul.)	(Nov.)		
Nyingchi	$286.5 \pm 106.6$	404.5 ±142.4	$257.8 \pm 123.0$	428.2 ±169.5	318.3±71.6	$2.47 \pm 0.7$
	(6:00)	(10:00)	(Jul.)	(Nov.)		
Qamdo	$738.4 \pm 517.7$	$1326.5 \pm 1048.8$	$469.3 \pm 217.7$	1887.1 ± 1132.0	766.1 ± 413.1	$-17.0 \pm 4.37$
	(5:00)	(21:00)	(Jul.)	(Dec.)		
Diqing	$455.8 \pm 257.8$	$798.7 \pm 427.1$	$419.0 \pm 221.2$	$793.2 \pm 666.8$	$531.4 \pm 156.8$	$-5.09 \pm 2.19$
	(5:00)	(10:00)	(Jul.)	(May)		

Table 3. Statistical summary of surface CO VMR in eight cities over the HTP region. All cities are organised as a function of increasing longitude.