

Interactive comment on “Local and Remote Response of Ozone to Arctic Stratospheric Circulation Extremes” by Hao-Jhe Hong and Thomas Reichler

Anonymous Referee #1

Received and published: 16 September 2020

The paper presents an analysis of stratospheric ozone anomalies associated with sudden stratospheric warmings (SSW), vortex intensification events (VI), as well as FW events at the end of the corresponding winters. MERRA-2 data is used and both the Arctic and the Tropics are examined. The transport mechanisms are examined using the Transformed Eulerian Mean formalism. The paper is very well written, the methods are valid and the interpretation of the results is correct. I only have minor comments that should be addressed before publication.

Minor comments

• **General:** *I find interesting the approach of interpolating the time axis in order to get common composited times for the SSW/VI event and for the FW. It helps bring out the outstanding ozone feature during the FW following VI events. However, I am not convinced by the terminology “the event’s duration” referring to the lapse time between the event’s central date and the FW. The final warming terminates the winter season, and while this indeed terminates the VI events, the SSW ends when the vortex recovers, not when it breaks down.*

Yes, we agree that “event duration” can be misleading. In the revised manuscript we now use “period” to better describe the time between the central date and the FW.

• **L134-138:** *I understand that S does not provide information on the photochemical changes as it is obtained as a residual and there are likely important numerical errors that prevent closing the budget, especially in a reanalysis system, where assimilation increments are included. However, you could check if the expected behavior is found in the S anomalies when referring to changes in photochemistry, e.g. for VI events in the polar region (L252-254) and in the tropics (L325-326).*

We performed the suggested analysis. Over the Arctic (Fig. R1) and in the upper stratosphere (above 10 hPa), the source term S is mostly anticorrelated with temperature T , as expected. Interestingly, during VIs and over the Arctic, there is some indication for chemical ozone depletion, consistent with previous studies (Isaksen et al., 2012; Manney et al., 2011, 2020). However, S over the Tropics (Fig. R2) does not follow this pattern and largely opposes the tendencies due to the vertical advection (Fig. 5g and 5h). We now mention this in the manuscript at line 255:

L255: *“We also examined the source term S (not shown) and find negative tendencies in the lower stratosphere (10 – 100 hPa) during and after the onset of VIs, indicative for temperature-driven heterogeneous ozone depletion as suggested by previous studies (Isaksen et al., 2012; Manney et al., 2011, 2020). In the upper stratosphere, S is as expected mostly anticorrelated with T .”*

Reference

Manney, G. L., Livesey, N. J., Santee, M. L., Froidevaux, L., Lambert, A., Lawrence, Z. D., Millán, L. F., Neu, J. L., Read, W. G., Schwartz, M. J., Fuller, R. A.: Record-low Arctic stratospheric ozone in 2020: MLS observations of chemical processes and comparisons with previous extreme winters, *Geophys. Res. Lett.*, 47, e2020GL089063, <https://doi.org/10.1029/2020GL089063>.

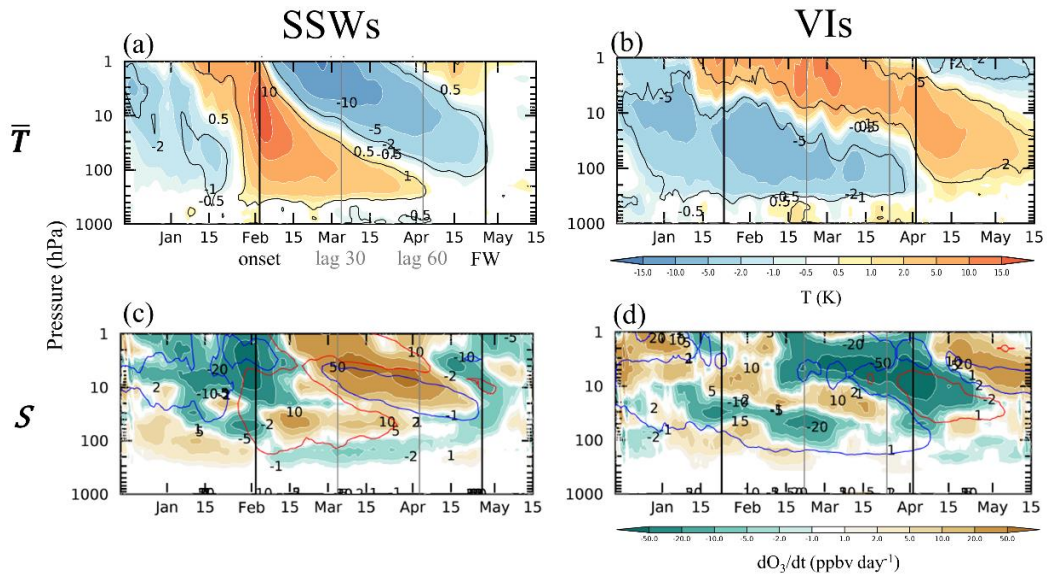


Figure R1. Composite anomalies for (left) SSWs and (right) VIs over the Arctic. Shown are time-height cross-sections for (a-b) temperature (K) (65°N-90°N) and (c-d) ozone tendency associated with S (ppbv day⁻¹).

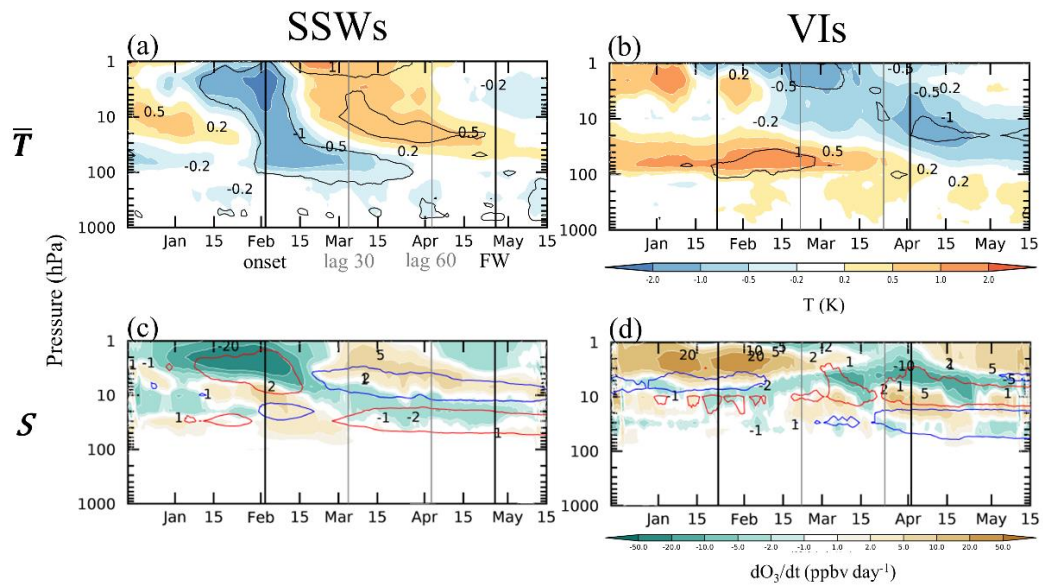


Figure R2. As Fig. R1, except for the Tropics ($\pm 15^\circ$).

• *Sections 3.1 and 4.1. Several of the features described in these sections have been previously shown in the article of de la Cámara et al. 2018 JGR (<https://doi.org/10.1002/2017JD028007>), such as the tropical upwelling or the wave activity for SSW composites, in both reanalysis and model data.*

We thank the reviewer for pointing this out. We now reference the paper to support its results.

• **L157: “one-tailed t-Student test”.** *This should be a two-tailed test, since the sample anomalies could be overestimating or underestimating the population anomalies (i.e. the null hypothesis is $\mu = 0$, not $\mu \leq 0$). This is important since for a 95% confidence level for the 8 VI events, $t_{0.025,7} = 2.365$ should be used instead of $t_{0.050,7} = 1.895$.*

We agree with the reviewer. In the revised manuscript, we now use for the figures a two-tailed t-test at the 95% confidence level. Note that our results suggest that the anomalies of temperature and ozone during VIs are still significant after the change.

• **L241-242:** *The small role of photochemical effects hypothesized here against the findings of Sagi et al. (2017) is consistent with the photochemical term shown in de la Cámara et al. (2018) ACP. Note that this paper does show the transport and chemistry relative contributions referred to L353-355 in a CCM, and that these CCM transport results are overall consistent with your reanalysis results.*

We also compared our ozone tendency results with de la Cámara et al. (2018). Although we use a different vertical system, the two results are quite consistent with each other in most of the stratosphere for SSW events. We followed your advice and changed our paper in two locations:

L241: *“Overall, this indicates that the decrease in mid-stratospheric ozone after SSWs is mainly of dynamical origin, consistent with de la Cámara et al. (2018). We note that this does not support the ideas of Sagi et al. (2017), who argue that the ozone decrease is due to chemical reactions involving NO_x species.”*

L353: *“Nevertheless, it would be interesting to evaluate the relative contributions from the dynamics and the chemistry in changing ozone during SSWs and VIs, using output from a range of coupled chemistry climate models (CCMs), similar in spirit to de la Cámara (2018) for SSWs using the WACCM model.”*

• **L304:** *These values are much lower than their Arctic counterparts, what is the relative TOC change?*

Fig. R5 shows the SSW and VI composites for column ozone as in Fig. 2 and Fig. 5. The blue lines in Fig. R5 are the percent column ozone anomaly with respect to climatology. The result suggests that changes in column ozone amount to 10%-12% of climatological values over the Arctic, while the ozone anomaly over the Tropics amount to only 0.5%-1%.

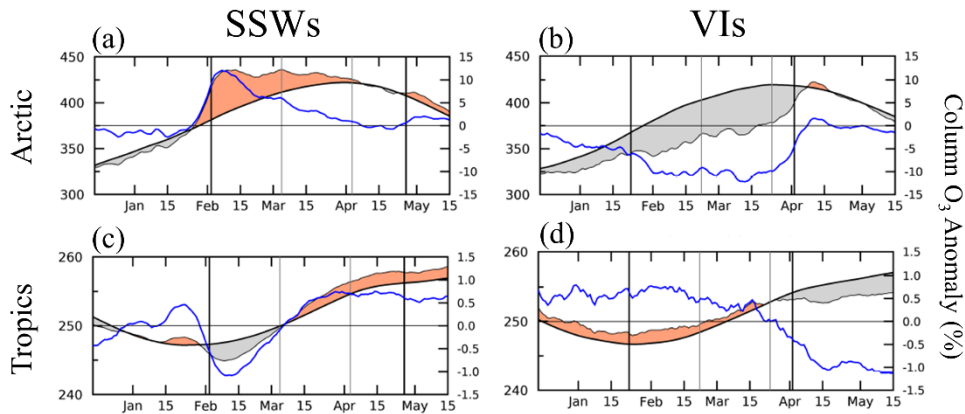


Figure R5. Column ozone composites (left) SSWs and (right) VIs. Blue line (right axis) represents the percentage of column ozone anomaly with respect to climatology.

We now add this information to our paper by adding/modifying the following sentences:

L301: “... during SSWs and VIs. The variations in tropical column ozone are rather small and amount to only $\sim 0.5 - 1\%$ of the climatological values, which can be compared to the $10 - 15\%$ changes seen over the Arctic. Nevertheless, the changes in tropical ozone are quite coherent and persistent. For SSWs, ...”

L303-305: “SSWs are followed by a small reduction in tropical column ozone by ~ 2.5 DU ($\sim -1\%$) and an increase by $\sim 1-2$ DU ($\sim 0.5\%$) after mid-March, which persists until late spring.”

L310: “During VIs (Fig. 5b), there are small tropical column ozone anomalies, which are mostly positive (~ 1 DU or 0.5%) and only become negative (~ 2 DU or 1%) after the FW.”

In the revised manuscript, we also replaced the column ozone tendency line (green lines in Fig. 2a-2b and Fig. 5a-5b) with the percentage of ozone anomalies (as in Fig. R5).

Technical

- **L35 onward:** Consider changing “transports” to “transport” throughout the paper?

This has been corrected.

- **L125:** change “p-coordinates” to “pressure coordinates”

This has been corrected.

- **L341:** “spectacular” Perhaps a more scientific term could be used (sudden/abrupt)?

We now use the word “remarkable” instead.

Interactive comment on “Local and Remote Response of Ozone to Arctic Stratospheric Circulation Extremes” by Hao-Jhe Hong and Thomas Reichler

Anonymous Referee #2

Received and published: 22 September 2020

This study analyzes ozone anomalies associated with stratospheric sudden warming, vortex intensification and final warming events based on MERRA-2 reanalysis data. Long-lasting anomalies are found both in over the Arctic and in the tropics following these extreme events. In particular, the ozone anomalies only become apparent after the QBO-related signals are removed. It is a useful exercise to document the evolution and distribution of ozone anomalies following the stratospheric circulation extremes. The paper is logically organized and clearly written. I have a few comments regarding some of the methodology and results. I recommend publication of the paper after these comments are addressed.

1. Removal of the QBO signals. It is not surprising that the SSW-related ozone anomalies are masked by the QBO-related anomalies, but it is somewhat surprising that authors tried several methods to filter out QBO and only one worked. What about the linear regression with a QBO index such as in Randel and Wu (2015 JAS)? The description of the method the authors chose (“subtract the preexisting ozone anomalies of each event from its subsequent daily ozone fields”) is not clear to me. The authors cited Gomez-Ecolar et al. (2014) for the method. But what described here does not seem to agree with any of the three methods described in Gomez-Ecolar et al. (2014). It sounds like calculating the difference of ozone between different periods. Then the resulting is actually ozone tendency rather than ozone anomalies itself.

We did not use linear regression in our analysis. Instead, we tested a similar method, taking into account the mean QBO ozone structure (Fig. 4a) and QBO phase during each event. However, we were not satisfied with the result and therefore used another filtering method.

Our filtering method defines a pre-existing QBO ozone signal from the mean ozone anomalies over day -60 to day -30 with respect to the SSW/VI central date. We subtract this QBO signal from the ozone anomalies associated with each circulation event before taking composites. This method assumes that the QBO time scale is much longer than the time scale of SSWs or VIs. Our method is similar to Gómez-Escolar et al. (2014) (Fig. 7 in their study) and Kodera (2006), however we defined a somewhat different time period for removing the QBO signal than the previous two studies.

To make this clearer, we reworded our manuscript as follows:

L291-299: “...by Dunkerton et al. (1988). To filter out the QBO influences from the tropical ozone, we define the QBO ozone signal as the mean ozone anomalies over day -60 to day -30 with respect to the SSW/VI central date, which is then subtracted

from the ozone associated with each Arctic circulation event. We use the resulting ozone anomalies for preparing Figs. 5c and 5d.”

2. Why there is a persistent minimum at about 20 hPa in the QBO-related ozone anomalies shown in Fig.4a? This feature seems unrealistic and is not seen in other studies (e.g. Fig. 1 of Tweedy et al. 2017 ACP).

Our result of the QBO ozone composite (Fig. 4a) is consistent with Tweedy et al. (2017) (Fig. R6), which only shows ozone anomalies between 10 and 70 hPa (right axis of their figure). The QBO ozone composite of Tweedy et al. (2017) (Fig. R6a) reveals a nodal point for ozone minimum between 10 and 20 hPa that is also seen in our result of Fig. 4a. As requested, we show below in Fig. R7 (black line) the evolution of ozone at 20 hPa, suggesting that the 20 hPa ozone undergoes a sign change within the QBO cycle. Another reason for the apparent discrepancies is that Tweedy et al. show the anomalies in percent, whereas we show them as absolute anomalies (in ppmv).

To clarify the similarity between Tweedy et al. (2017) and our result, we now reword our manuscript as follows:

L283: “Our results (Fig. 4a) are in good agreement with their study, e.g., there is a nodal point of small ozone variations between 10 and 20 hPa, with much stronger variations above and below.”

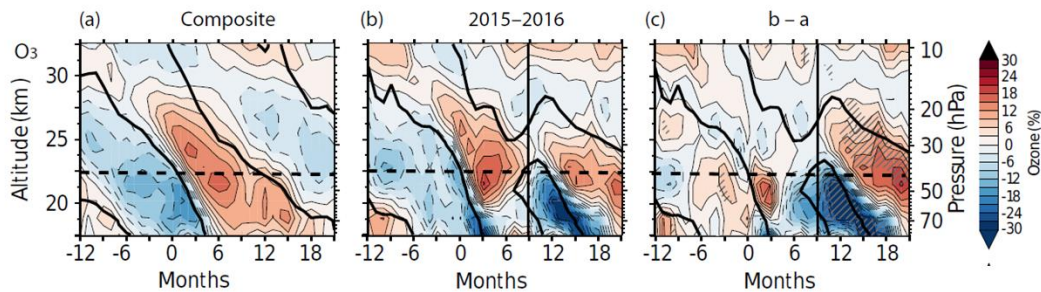


Figure R6. Time-height cross-sections for QBO ozone. Shown are ozone anomalies for (a) QBO composite, (b) 2015-2016 QBO event and (c) b-a. Adapted from Tweedy et al. (2017), Fig. 1.

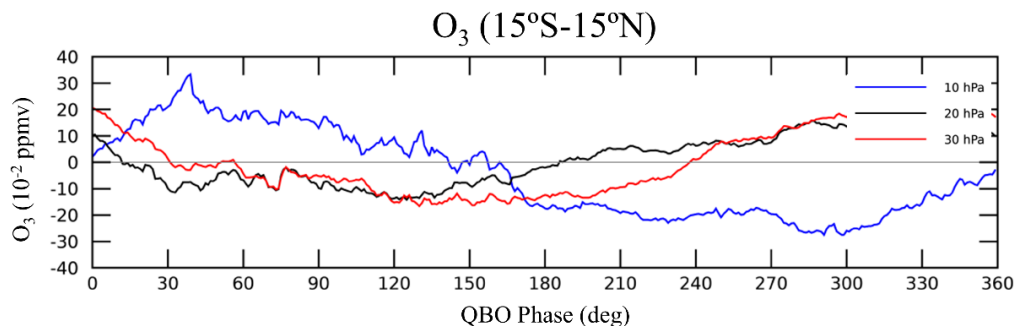


Figure R7. Composite tropical ozone anomalies for QBO events. Shown are anomalies at (blue) 10 hPa, (black) 20 hPa, and (red) 30 hPa.

3. The authors show the ozone tendency due to residual mean circulation and eddy flux convergence, and the eddy convergence term contribute significantly over the Arctic. Can the authors elaborate a bit more on the physical process associated with the eddy convergence, especially what determines the sign of this term?

To answer your question, we revised our manuscript and explained in more detail the meaning of the eddy flux convergence term in Sect. 2.3:

L132: *“The eddy flux convergence contains effects that are not explained by the advection of zonal mean ozone by the zonal mean circulation. The convergence is associated with transports of zonal disturbances in ozone by zonal disturbances in meridional or vertical velocity. In the stratosphere, these disturbances (or eddies) are primarily due to upward propagating planetary waves. The convergence term indicates that covariance between eddy velocities and ozone can transport ozone, and that where this eddy ozone flux converges a zonal mean ozone tendency can be induced. For example, a northward ozone flux is created if the signs of the meridional velocity and the ozone perturbations tend to be the same, and if this flux decreases in the northward direction (converges), it would create a positive ozone tendency in the zonal mean. Our result (not shown) suggests that the meridional component of the eddy flux convergence (the first term of the \mathbf{M} -vector in equation (1) dominates the vertical component over most of the stratosphere.”*

4. Line 118: “180-day smoothed” Do the authors mean a running mean with 180 day window?

Yes, we understand our wording is not clear. We now revise our manuscript as follows:

L118: *“We use a 180-day running mean window to smooth the zonal-mean equatorial ($\pm 5^\circ$) zonal wind at 30 hPa (UEQ30) and determine the phase of the QBO.”*

5. Line 208-209: Why is the magnitude of the ozone anomalies associated with SSW differ so much between Hocke et al. (2015) and this study?

The differences are only apparent. While we show area-weighted latitudinal averages over 65°N - 90°N (with a maximum of ~ 50 DU), Fig. 2 in Hocke et al. (2015) shows larger ozone anomalies (up to 90 DU) only close to the pole (which represents a small area). Therefore, we believe that our result is consistent with Hocke et al. (2015).

In the revised manuscript, we add a clarification at line 209 as follows:

L209: *“However, we note that the differences are only apparent, as we show area-weighted latitudinal averages of column ozone and as the extreme ozone increases in Hocke et al. (2015) occur only close to the pole.”*

6. Line 225: From Fig. 2, the FW anomalies following a VI event is stronger and extends to the lower stratosphere.

Yes, actually there are some negative ozone anomalies in the lower stratosphere at the FW following the VIs. We revised our manuscript as follows:

L224-226: *“The structure of these anomalies is somewhat similar to that of SSWs, except that they are weakly negative in the lowermost stratosphere.”*

Interactive comment on “Local and Remote Response of Ozone to Arctic Stratospheric Circulation Extremes” by Hao-Jhe Hong and Thomas Reichler

Anonymous Referee #3

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This manuscript shows the dynamical features of SSW and VIs from the boreal early winter to late spring in both region, Arctic and Tropics. The present study found new aspect on the dynamical impact of the final warming at the case of Vis winter. Further, there are descriptions of ozone fields at Arctic and tropics on SSW and Vis winters. The manuscript was well written, the present manuscript will be published after modifying some minor corrections and/or answer to the reviewer comments.

Minor comments:

Abstract: *It is interesting points that should be described in the Abstract, that is the quantitative discussion of ozone change (Figs.2 (a,b) and Figs.5 (a,b)). The ozone change was large at both case of SSW and Vis with FW in the Arctic, on the other hand, the ozone change was small but same amount for both case in the tropics.*

We revised the abstract by adding specific numbers for the changes in column ozone in the following way:

L15-17: *“Over the Arctic and during sudden warmings, ozone undergoes a rapid and long-lasting increase of up to ~50 DU, which only gradually.... In contrast, vortex intensifications are passive events, associated with gradual decreases in Arctic ozone that reach ~40 DU during late winter and decay thereafter.”*

L20-21: *“After controlling for this effect, small but coherent reductions in tropical ozone can be seen during the onset of sudden warmings (~2.5 DU), and also during the final warmings that follow vortex intensifications (~2 DU).”*

Line 41 “become easterly”: *The major warming event accepts the reversal of zonal wind direction from westerly to easterly, however the warming event does not always reverse the wind direction, like for the minor warming.*

We now revised the manuscript as follows:

L37-38: *“At times, the bursts of waves and their interaction with the polar vortex are strong enough to create so-called major Stratospheric Sudden Warming Events (SSWs)...”*

Line 253-255 “In contrast, the negative anomalies. . .”: *Figs.2 (c, d) : Do Figs.2 (c,d) show the anomaly from climatology? If so, the authors should add the description of the anomaly from what.*

Yes, Figs. 2c and 2d are daily ozone anomalies, and we corrected the caption of figure 2 to make this clearer:

Caption 2: *“Figure 2. Arctic ozone Remaining panels are anomalous time-height cross-sections of ...”*.

The calculation of daily ozone anomaly is already described in section 2.1, where we write:

L96-99: *“We compute daily climatologies from MERRA-2 by averaging each day of the year over the entire record and smoothing over the seasonal cycle using 10-day running means. Daily anomalies are obtained by subtracting the climatologies from the daily data.”*

Line 264 “the variations of w (Fig.3a)”: If the QBO variation remains in the residual vertical velocity and temperature fields, the authors should note the fact.

Yes, the QBO variations are still contained in the diagnostic of the dynamical quantities. To make this clearer, we now add one sentence after line 263:

“Note that no filtering has been applied to this figure and that the shown changes are due to both the remote impacts from the Arctic circulation events and the local effects from events like the QBO.”

Line 278 Fig4b: The year of 2016 for Vis case is absent.

As described in section 2.2, we did not include the anomalous 2015-2016 QBO event in our analysis. We revised the caption of Fig. 4 as follows:

Caption 4: *“Figure 4. Composites for QBO events ...; S and V on the right axis is the mean UEQ30 of all SSWs and all VIs (except 2016), respectively. The 2015-2016 QBO event has been purposefully excluded from this analysis due to the anomalous nature of this event.*

Interactive comment on “Local and Remote Response of Ozone to Arctic Stratospheric Circulation Extremes” by Hao-Jhe Hong and Thomas Reichler

Author Comment

Dear reviewers,

We also apply the following changes in the revised manuscript to better describe the work of Lubis et al. (2017):

L355-358: "We also did not explicitly consider so-called Downward planetary Wave Coupling events (DWCs) (Lubis et al., 2017), relatively short-lived events (< 10 days) associated with increases in ozone before and decreases during the event, leading to a relatively small net response. Our VI events also need to be distinguished from so-called reflective winters, introduced by Shaw and Perlwitz (2013) and discussed by Lubis et al. (2017) to indicate winters in which wave reflection dominates. Although defined in different ways, there is some overlap between years with VIs and reflective winters and they are both associated with negative anomalies in wave driving and ozone."

Local and Remote Response of Ozone to Arctic Stratospheric Circulation Extremes

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Abstract.

Intense natural circulation variability associated with stratospheric sudden warmings, vortex intensifications, and final warmings is a typical feature of the winter Arctic stratosphere. The attendant changes in transport, mixing, and temperature create pronounced perturbations in stratospheric ozone. Understanding these perturbations is important because of their potential feedbacks with the circulation and because ozone is a key trace gas of the stratosphere. Here, we use MERRA-2 reanalysis to contrast the typical spatiotemporal structure of ozone during sudden warming and vortex intensification events. We examine the changes of ozone in both the Arctic and the Tropics, document the underlying dynamical mechanisms for the observed changes, and analyze the entire life-cycle of the stratospheric events - from the event onset in mid-winter to the final warming in early spring. Over the Arctic and during sudden warmings, ozone undergoes a rapid and long-lasting increase of up to ~50 DU, which only gradually decays to climatology before the final warming. In contrast, vortex intensifications are passive events, associated with gradual decreases in Arctic ozone that reach ~40 DU during late winter and decay thereafter. The persistent loss of Arctic ozone during vortex intensifications is dramatically compensated by sudden-warming-like increases after the final warming. In the Tropics, the changes in ozone from Arctic circulation events are obscured by the influences from the quasi-biennial oscillation. After controlling for this effect, small but coherent reductions in tropical ozone can be seen during the onset of sudden warmings (~2.5 DU), and also during the final warmings that follow vortex intensifications (~2 DU). Our results demonstrate that Arctic circulation extremes have significant local and remote influences on the distribution of stratospheric ozone.

1 Introduction

25 The wintertime Arctic stratosphere is characterized by a number of dynamical, chemical, and physical processes that are coupled to each other in intriguing ways. For example, extreme stratospheric circulation events from the interaction (or lack thereof) of upward propagating planetary-scale Rossby waves with the polar vortex create a pronounced dynamical variability in the Arctic. Large concentration of ozone is another important characteristic of the Arctic stratosphere. Ozone is an effective absorber for solar radiation and an important player in the coupling between the chemistry, radiation, and

30 dynamics. The diabatic heating from ozone impacts the temperatures and the winds, and the induced dynamical transports and photochemical reactions again impact the ozone. The feedback between ozone and the circulation may sustain the circulation anomalies and modify the stratospheric sensitivity to external forcings (Hartmann et al., 2000). Ozone is also important for the protection of life on Earth by absorbing harmful ultraviolet radiation. Taken together, ozone is a crucial stratospheric constituent, and understanding the factors that influence its distribution is a critical goal of climate research.

35 Ozone in the Arctic lower stratosphere is mostly controlled by transports. The [transport intensifies](#) in the winter hemisphere (Randel, 1993; Randel et al., 2002) to create a springtime total ozone maximum at high latitudes. The seasonality of the transports is associated with an intensification of the upward propagating Rossby waves in winter. At times, the bursts of waves and their interaction with the polar vortex are strong enough to create so-called [major](#) Stratospheric Sudden Warming Events (SSWs) (McIntyre, 1982; Limpasuvan et al., 2004; Polvani and Waugh, 2004), arguably the most important
40 form of stratospheric circulation events. In the process, polar temperatures increase rapidly, reverse the climatological equator-to-pole temperature gradient, and cause the normal westerly flow of the vortex to become easterly (Scherhag, 1952). SSWs occur in about two of every three years (Butler et al., 2017), most often in January or February (Horan and Reichler, 2017).

Past studies pointed out the close coupling between the stratospheric dynamics and Arctic ozone (e.g., Leovy et al., 1985; Ma et al., 2004), with a positive correlation between polar ozone tendencies and the stratospheric wave driving (Randel et al., 2002). The coupling leads to enhanced poleward ozone transports during SSWs and creates persistent ozone anomalies in the lowermost stratosphere (Butler et al., 2017; Hocke et al., 2015). De la Cámara et al. (2018**b**) showed that the initial increase in ozone after SSWs is mainly driven by isentropic eddy fluxes associated with the enhanced wave driving, while the subsequent recovery of ozone can be attributed to the competing effects between cross-isentropic advection and
50 irreversible isentropic mixing.

It is perhaps less well-known that the influence of SSWs on ozone can also influence the Tropics. Randel (1993) demonstrated how vertical transports from the 1979/80 SSW affected tropical ozone in the lower stratosphere and how the changes in ozone were correlated with temperatures in the upper stratosphere. The SSW-related influences on the Tropics also imprint on the variability of temperature and water vapor there (Gómez-Escolar et al., 2014; Tao et al., 2015). However,
55 the SSW effect on tropical ozone is superimposed on the effects from the Quasi-Biennial Oscillation (QBO), downward propagating westerly and easterly zonal wind anomalies with a cycle of ~28 months (Baldwin et al., 2001; Coy et al., 2016; Randel and Wu, 1996) that also influence ozone.

The winter Arctic stratosphere not only witnesses occasional SSWs. A sustained lack of stratospheric wave driving can create the opposite events to SSWs, so-called Vortex Intensification events (VIs). VIs are characterized by an unusually
60 strong and cold polar vortex (Limpasuvan et al., 2005), and reduced transports of ozone into the pole region (Isaksen et al., 2012). The extreme cold during VIs favors halogen-induced chemical ozone depletion, which, in combination with the weakened transport, leads to record low levels of ozone that can be comparable in magnitude to its southern hemispheric counterpart (Isaksen et al., 2012; Manney et al., 2011). A good example is the most recent winter 2019/20, which

experienced an exceptionally strong, cold, and persistent Arctic stratospheric polar vortex, and which led to record-breaking
65 Arctic ozone depletion.

Another important class of stratospheric circulation events are stratospheric Final Warming events (FWs). FWs occur every year at the end of winter, representing the final breakdown of the polar vortex due to the seasonal increase in solar heating. FWs are often triggered by pulses of increased wave activity and can be considered as SSWs that conclude the winter season (Black et al., 2007). There also exists an interesting temporal relationship between FWs, SSWs, and VIs: FWs
70 that are preceded by SSWs in the same winter tend to occur significantly later than the mean FW date (~mid-April, Horan and Reichler, 2017), and FWs that are preceded by non-SSW winters (i.e., neutral winter and VIs) tend to be relatively early (Hu et al., 2014). This can be explained from the delayed relationship between vortex strength and wave driving. An SSW, for example, is usually followed by reduced wave activity and hence a stronger vortex, which then breaks down later in spring. The changes in FW timing also impact the levels of Arctic ozone: Manney and Lawrence (2016) showed that the
75 chemical ozone loss from the 2016 VI was disrupted by an early FW at the beginning of March and suggested that FWs may have comparable effects on Arctic ozone as SSWs.

While the aforementioned studies have started to investigate the response of ozone in the Arctic to SSWs, the response of ozone in the Tropics and also to VI and FW events has received little attention so far. This study intends to fill this gap and refine the existing knowledge about the spatiotemporal relationship between ozone and a range of Arctic stratospheric
80 circulation events using a modern observation-based perspective. We achieve this by taking a comparative approach that contrasts the often-opposing ozone behavior between SSWs and VIs, and between the Arctic and the Tropics. Time is another distinctive aspect of this study, as we cover the entire life-cycle of the stratospheric circulation events from the event onset in the middle of winter to the date of the FW at the end of winter. We also clarify the role of the associated dynamical and photochemical processes in changing ozone. Overall, our goal is to provide an up-to-date observation-based view of the
85 global natural dynamics-driven variability of stratospheric ozone. This is not only of interest in its own right but also provides an observational baseline for ozone behavior during stratospheric circulation events that can be used for the validation of coupled chemistry-climate models.

This paper is structured as follows. In Sect. 2, we describe the data and methods used in this study. In Sect. 3, we demonstrate the ozone response in the Arctic, while in Sect. 4 we continue our discussion for the Tropics. A summary and
90 conclusion are provided in Sect. 5.

2 Data and Methods

2.1 MERRA-2 Data

We use 1980-2018 daily fields from the MERRA-2 reanalysis (Bosilovich et al., 2015) at a horizontal resolution of 1.5° and 37 levels ranging from 1000 to 0.1 hPa. MERRA-2 also provides ozone, which is based on retrievals from the SBUV
95 (January 1980-September 2004) and Aura MLS/OMI (October 2004-present) instruments (Davis et al., 2017) and on a

simple ozone scheme (Rienecker et al., 2008). MERRA-2 has been shown to perform well for ozone through much of the stratosphere (Davis et al., 2017; Wargan et al., 2017). Most of our calculations are based on zonal mean quantities. We compute daily climatologies from MERRA-2 by averaging each day of the year over the entire record and smoothing over the seasonal cycle using 10-day running means. Daily anomalies are obtained by subtracting the climatologies from the daily data.

2.2 Event Definition

In defining SSWs and FWs, we follow the widely used prescription by Charlton and Polvani (2007). An SSW is detected when the zonal-mean zonal wind at 10 hPa and 60° N (U1060) switches from westerly to easterly (the central date of the event) during November-March and returns to westerly for at least ten consecutive days before 30 April. If the return to westerly condition is not fulfilled, the event is considered as the FW of the year. Two or more SSWs in the same winter must be separated by consecutive westerlies for at least 20 days. Since we are interested in the evolution of ozone over the life-cycle of SSWs from the middle to the end of the winter, we only consider mid-winter SSWs during January or February. We also discard mid-winter SSW events that are followed by another, potentially disturbing, SSW, leading to the exclusion of only one event.

Our definition of mid-winter VIs is also based on U1060, but we first low-pass filter the data using 20-day running means. A mid-winter VI occurs when the smoothed daily U1060 anomaly during January or February exceeds one standard deviation (16 m s⁻¹), marking the central date of the VI. Like SSWs, two VIs in the same winter must be separated by at least 20 days. We only consider VIs that are not followed by another VI or SSW.

As shown in Table 1, our definitions lead to 15 SSWs and 8 VIs. For SSWs, the mean central date and the associated FW date are 3 February and 26 April, respectively, leading to a mean period of 83 days (ranging from 54 to 117). VIs have a mean central date on 23 January and an associated FW date on 2 April. This translates into a mean period of 70 days (ranging from 44 to 91). Note that SSWs are longer in period than VIs, consistent with the findings by Hu et al. (2014) that SSW winters are associated with FW dates that are on average late compared to the climatological mean FW date.

We use a 180-day running mean window to smooth the zonal-mean equatorial ($\pm 5^\circ$) zonal wind at 30 hPa (UEQ30) and determine the phase of the QBO. A QBO cycle is defined as the period between two consecutive positive UEQ30 maxima, and the UEQ30 minimum in between is considered as the midpoint of the cycle. We exclude the anomalous QBO cycle of 2015-2016 (Newman et al., 2016) from our analysis and obtain 16 QBO cycles over the 1980-2018 period.

2.3 Ozone and Dynamics Diagnostics

The changes in zonal-mean ozone ($\bar{\chi}$) are investigated using the transformed Eulerian mean (TEM) approach. Following Andrews et al. (1987), the TEM tracer transport equation in pressure coordinates

$$\bar{\chi}_t = -\bar{v}^* \bar{\chi}_y - \bar{\omega}^* \bar{\chi}_p - \rho_0^{-1} \nabla \cdot \mathbf{M} + \bar{S}, \quad (1)$$

is used to decompose the ozone tendency ($\overline{\chi_t}$) into two advection terms associated with the residual mean circulation, one term due to eddy flux convergence ($-\rho_0^{-1}\nabla \cdot \mathbf{M}$), and a source term (\overline{S}) that represents the effects of chemistry on ozone.

130 ~~The eddy flux convergence term represents the effects of the resolved eddies in transporting and mixing ozone (Andrews et al., 1987).~~ Here, $\overline{v^*}$ and $\overline{\omega^*}$ are the components of the residual mean circulation, ρ_0 is the basic state density, \mathbf{M} is an eddy flux vector given by

$$\mathbf{M} = [\rho_0(\overline{v'\chi'} - \overline{v'\theta'} \overline{\chi_p}/\overline{\theta_p}), \rho_0(\overline{\omega'\chi'} + \overline{v'\theta'} \overline{\chi_y}/\overline{\theta_p})],$$

135 where overbars denote zonal means, primes are deviations from zonal means, and the other terms are standard notation. ~~The eddy flux convergence contains effects that are not explained by the advection of zonal mean ozone by the zonal mean circulation. The convergence is associated with transports of zonal disturbances in ozone by zonal disturbances in meridional or vertical velocity. In the stratosphere, these disturbances (or eddies) are primarily due to upward propagating planetary waves. The convergence term indicates that covariance between eddy velocities and ozone can transport ozone, and that where this eddy ozone flux converges a zonal mean ozone tendency can be induced. For example, a northward ozone flux is created if the signs of the meridional velocity and the ozone perturbations tend to be the same, and if this flux decreases in the northward direction (converges), it would create a positive ozone tendency in the zonal mean. Our result (not shown) suggests that the meridional component of the eddy flux convergence (the first term of the \mathbf{M} -vector in equation (1)) dominates the vertical component over most of the stratosphere.~~

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The ozone tendency $\overline{\chi_t}$ is calculated by taking forward differences in time of daily ozone, and the chemical source term \overline{S} is the residual between $\overline{\chi_t}$ and the sum of the three dynamical terms of Eq. (1). We note that the resulting \overline{S} does not exclusively reflect the chemical production or destruction of ozone because of unavoidable errors of MERRA-2 and computational uncertainties. For example, in the absence of observations, the MERRA-2 ozone is calculated from a simple parameterization (Rienecker et al., 2008), which can result in considerable errors. Because of this uncertainty, and also because of the focus of this study on the dynamical impacts, we do not show the \overline{S} term.

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We use F_p , the vertical component of the quasi-geostrophic Eliassen-Palm (EP) flux (Eliassen and Palm, 1961), to diagnose the upward propagating Rossby wave activity. Following Andrews et al. (1987), F_p is given by

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$$F_p = -a \cos \phi f \frac{\overline{v'\theta'}}{\overline{\theta_p}}, \quad (2)$$

where all symbols are standard notation. In our analysis, we reverse the sign of F_p so that positive F_p corresponds to upward propagation. We focus on F_p at 100 hPa averaged over 40° N-80° N and refer to this quantity as the stratospheric wave driving.

155 2.4 Event Compositing

Traditional composites take averages of various events centered on specific dates (e.g., Butler et al., 2017). However, in the present study, we are interested in the behavior of ozone during the entire life cycle of stratospheric circulation events, with each event having a somewhat different period. The life cycle starts at the central date and ends with the FW at the end of

winter. Since the period (i.e., the time between the central date and the FW) differs from event to event, we somewhat
 160 modify the traditional compositing technique. Our approach is based on the mean central date of all selected SSWs (or VIs)
 (\bar{t}_0) and the mean date of their associated FWs ($\overline{t_{FW}}$). We then use linear interpolation in time to align the dates of the
 individual events (t_0, t_{FW}) with the composite mean dates ($\bar{t}_0, \overline{t_{FW}}$). Mathematically, this can be written as

$$\bar{t} = \bar{t}_0 + (t - t_0) * \frac{\overline{t_{FW}} - \bar{t}_0}{t_{FW} - t_0}, \quad (3)$$

where \bar{t} denotes the time of the composite and t the time of individual events. The interpolation can be interpreted as a
 165 stretching or squishing of the time axis so that all data during t_0 (t_{FW}) are aligned with \bar{t}_0 ($\overline{t_{FW}}$). The mean period ($\overline{t_{FW}} - \bar{t}_0$)
 of SSWs (VIs) is then 83 (70) days. The period of individual events ($t_{FW} - t_0$) is shown in Table 1. We use this technique to
 create composites of various quantities at daily intervals. A two-tailed Student's t-test at the 95% confidence level is used to
 test the statistical significance of the composite mean anomalies against the null hypothesis of zero anomalies.

3 Arctic Ozone

170 3.1 Arctic Circulation Changes

We begin our discussion of how Arctic ozone evolves during SSWs and VIs by presenting some key dynamical quantities,
 which will then guide the interpretation of our subsequent results. Fig. 1 shows the evolution of composite anomalies in the
 stratospheric wave driving (top), the vertical component of the residual circulation (middle), and temperature (bottom) over
 the life cycle of SSWs (left) and VIs (right).

175 SSWs (Fig. 1, left) are typically preceded by enhanced stratospheric wave driving, starting at a negative lag of ~15 days
 (Fig. 1a). This leads to the breakdown of the polar vortex and marks the onset of the SSW (Limpasuvan et al., 2004). After
 the onset, the wave driving decreases rapidly and becomes negative, contributing to the over-recovery of the vortex in the
 upper stratosphere, reminiscent of so-called polar-night jet events ([de la Cámara et al., 2018a](#); [Hitchcock and Shepherd,
 2013](#); [Kuroda and Kodera, 2001](#); ~~[Hitchcock and Shepherd, 2013](#)~~). As pointed out by Plumb and Eluszkiewicz (1999) and
 180 demonstrated by Fig. 1c, this cyclic nature of the wave driving imprints on the residual circulation of the entire stratosphere.
 Fig. 1c shows that the vertical component of the residual circulation ($\bar{\omega}^*$) over the Arctic varies consistently with the wave
 driving, with enhanced downwelling during onset (reddish colors), followed by a long period of enhanced upwelling (bluish
 colors). The cycle ends at the end of winter, with somewhat enhanced wave driving and subsequent downwelling during the
 FW. Arctic temperatures (Fig. 1e) are characterized by cooling before the SSW, strong warming in the middle to lower
 185 stratosphere during and after the onset, and cooling after the onset in the upper to the middle stratosphere. The patterns of
 warming and cooling following the onset give the impression of a downward propagation. However, the cooling in the upper
 stratosphere is associated with the aforementioned suppressed wave driving and subsequent radiative cooling ([Hitchcock and
 Shepherd, 2013](#); [Limpasuvan et al., 2004](#)), and the persistence of the warming in the lower stratosphere is related to the long
 radiative time scale in this part of the stratosphere.

190 VIs (Fig. 1, right) are in many respects opposite to SSWs. As explained in Limpasuvan et al. (2005), VIs evolve relatively slowly and result from the sustained lack of stratospheric wave driving, leading to the gradual strengthening and cooling of the vortex. As shown by Fig. 1b, the wave driving is anomalously small, starting several weeks before onset and minimizing at about one week after onset. This is different from SSWs, as the wave driving during SSWs changes much more abruptly during onset. Long after the onset of VIs, the wave driving increases again, first more intermittently, and then more
195 systematically during the FW. We note that the magnitude of the wave driving associated with the FW is quite large and comparable to that of SSWs during onset. This may be attributable to the sustained suppression of wave driving during VI onset, contributing to the enhanced release of wave activity after the event and a relatively early FW. Also, the relatively strong polar vortex after VIs (not shown) is conducive for upward propagating wave activity into the stratosphere.

As for SSWs, changes in the Arctic $\bar{\omega}^*$ during VIs (Fig. 1d) agree well with the evolution of wave driving. The upwelling
200 maximizes one week after VI onset, followed by a period of intermittent downwelling before the FW (see also Limpasuvan et al., 2005). VIs are also associated with pronounced and persistent Arctic cooling (Fig. 1f) in the lower stratosphere, which is in contrast to the significant warming that starts about one week before VI onset in the upper stratosphere. The warming slowly propagates downward, persists until spring, and finally becomes part of the FW that concludes the winter season. The timing and strength of the FW is another important difference between SSWs and VIs. While FWs after SSWs tend to be late
205 and mostly represent a transition into climatology, FWs after VIs occur early, are relatively strong, and contribute to a pronounced weakening and warming of the vortex.

3.2 Arctic Ozone Changes

The above-described dynamical perturbations are associated with significant changes in transport of stratospheric ozone and its temperature-dependent photochemical reaction rates. As has been shown to some extent before (Butler et al., 2017; de la
210 Cámara et al., 2018b; Hocke et al., 2015), and as we will show in more detail next, this has major consequences for the distribution of stratospheric ozone.

We first examine the composite evolution of Arctic column ozone (i.e., the vertically integrated ozone amount) during SSWs (Fig. 2a). Red and gray shading indicate the deviation of the column ozone from its climatology (thick black curve), and the green line shows the percent column ozone anomaly with respect to climatology. As noted by Randel et al. (2002),
215 the ozone tendency is well correlated with the wave driving (Fig. 1a). Before onset, there is a subtle decrease in column ozone, presumably related to the anomalously strong and cold vortex during this time (Fig. 1e) and the reduced ozone transport into the polar regions. Within the first 10 days following the SSW onset, the column ozone anomalies rapidly increase by ~50 DU and persist for up to 60 days until late winter. Hocke et al. (2015) suggested that the increases in column ozone after SSWs amount to up to 90 DU over the Arctic, which is nearly twice of what we find. However, we note that the
220 differences are only apparent, as we show area-weighted latitudinal averages of column ozone and as the extreme ozone increases in Hocke et al. (2015) occur only close to the pole. The vertically resolved Arctic ozone mixing ratio (Fig. 2c) shows a more complicated picture. There is a pronounced reduction in ozone in the middle and upper stratosphere after

SSWs, which seems to be slowly descending downward. This decrease in mid-stratospheric ozone, which starts about one month after SSWs, has also been noted by Sagi et al. (2017). Ozone in the upper stratosphere also undergoes a complicated evolution. The negative anomalies above 5 hPa exist only shortly during onset. They are followed by persistent positive anomalies, which again tend to descend downward by mid-March, diminish by April, and reemerge at mid-stratospheric levels by the end of April as a consequence of the FW.

Next, we examine the evolution of Arctic ozone during VIs (Fig. 2, right). Column ozone (Fig. 2b) is anomalously negative over the entire VI life cycle, minimizing at about -40 DU by mid-March. ~~As with SSWs, the anomalous tendency of column ozone (green line) resembles the wave driving (Fig. 1b) and is strongly positive during the FW.~~ Fig. 2d demonstrates that the negative ozone anomalies maximize in the middle stratosphere at ~10 days after onset and also tend to propagate downward into the lower stratosphere. These anomalies are particularly long-lasting in the lower stratosphere, where they exist for more than 60 days until the FW. This composite behavior is very similar to the case study by Manney and Lawrence (2016), who reported that the rapid Arctic chemical ozone loss during winter 2015/16 was abruptly terminated by the early FW in March. Ozone anomalies are also negative in the upper stratosphere, where they persist throughout the VI life-cycle and tend to descend after the FW. At the FW, there are strongly positive ozone anomalies in the middle stratosphere. The structure of these anomalies is somewhat similar to that of SSWs, except that they are weakly negative in the lowermost stratosphere.

We now explore the role of the dynamical mechanisms that create the changes in ozone. From the TEM tracer transport equation Eq. (1) it is clear that several processes are involved. Fig. 2e-2j present the total time tendencies of ozone (e-f) and the contributions to it from vertical advection (g-h) and eddy flux convergence (i-j). The horizontal advection term is generally small and therefore omitted. For better orientation, the red and blue contours reproduce a constant ozone mixing ratio anomaly from Figs. 2c and 2d.

The negative Arctic ozone anomalies in early winter before SSWs are partly the result of reduced eddy flux convergences (Fig. 2i) and vertical transports (Fig. 2g). The strong positive ozone tendencies close to the onset of SSWs, which are responsible for the increase in ozone after SSWs, result mainly from the convergence of eddy fluxes (Fig. 2i) (see also de la Cámara et al., 2018b), triggered by the enhanced wave driving associated with SSWs (Fig. 1a). The downward transport of ozone by the enhanced residual circulation also contributes to the positive tendencies during onset, in particular in the lower stratosphere (Fig. 2g). After SSWs, the suppressed planetary wave activity leads to a sustained reduction of eddy transports, and hence negative ozone tendencies in the middle and lower stratosphere. At the same time, the vertical advection of ozone is anomalously negative in the middle stratosphere after SSWs. Both effects lead to the gradual decay of the strongly positive ozone anomalies right after onset and eventually create the abovementioned banded structure of negative ozone in the middle stratosphere. Overall, this indicates that the decrease in mid-stratospheric ozone after SSWs is mainly of dynamical origin, consistent with de la Cámara et al. (2018b). We note that this does not support the ideas of Sagi et al. (2017), who argue that the ozone decrease is due to chemical reactions involving NO_x species. During the time of the FW, the eddy flux convergence becomes somewhat positive (Fig. 2i), overall leading to ozone mixing ratios that are close to climatology. In the

upper stratosphere, the temperature-dependent photochemistry plays a dominant role for ozone. There, ozone is mostly anti-correlated with temperature (Craig and Ohring, 1958), which can be seen by comparing Fig. 1e (for temperature) with Fig. 2c (for ozone).

260 The VI related total Arctic ozone tendencies (Fig. 2f) are mostly equal but opposite in sign to that of SSWs. VIs are passive events that develop gradually by radiative cooling out to space, and the related negative ozone anomalies appear long before the actual onset (Fig. 2d), related to periods of negative tendencies before and during VI onset (Fig. 2f). The tendencies are related to reduced eddy transports in the upper half (Fig. 2j) and reduced vertical advection in the lower half of the stratosphere (Fig. 2h). Ozone in the upper stratosphere slowly recovers towards climatology, mostly due to increases
265 in eddy transport associated with pulses of planetary waves that restore the vortex back to normal. However, the positive eddy transport is counteracted by the photochemical effect as the temperature is anomalously warm in this layer (Fig. 1f). In contrast, the negative ozone anomalies in the lower stratosphere are sustained by reduced vertical advection (Fig. 2h) until mid-March. We also examined the source term S (not shown) and find negative tendencies in the lower stratosphere (10 – 100 hPa) during and after the onset of VIs, indicative for temperature-driven heterogeneous ozone depletion as suggested by previous studies (Isaksen et al. 2012; Manney et al., 2011, 2020). In the upper stratosphere, S is as expected mostly anticorrelated with T . As explained before, FWs that follow VIs tend to be relatively strong and somewhat resemble SSWs, leading to sizeable increases in Arctic ozone. As with SSWs, this is associated with positive eddy transports in the upper half (Fig. 2j) and positive vertical advection in the lower half of the stratosphere (Fig. 2h). The two effects compensate for the prior ozone deficits, leading to an overall recovery of the column ozone anomalies (Fig. 2b).

275 4 Tropical Ozone

4.1 Tropical Circulation Changes

We now turn our attention to the Tropics, defined as the $\pm 15^\circ$ latitude band. Tropical ozone is changing in response to Arctic circulation events because of the global nature of the meridional overturning and its role in the transport of ozone (Randel, 1993). We start our discussion by focusing on the changing dynamics in the Tropics during Arctic circulation events (Fig. 3).
280 Note that no filtering has been applied to this figure and that the shown changes are due to both the remote impacts from the Arctic circulation events and the local effects from events like the QBO. During SSWs, the variations of $\bar{\omega}^*$ (Fig. 3a) are largely opposite to that in the Arctic (Fig. 1c) (de la Cámara et al., 2018a), except during the time of the FW. This demonstrates that the global nature of the enhanced residual circulation during SSWs also affects the Tropics, leading to stronger upwelling and cooling. The cooling persists in the lower stratosphere, but quickly transitions into warming in the
285 middle and upper stratosphere (Fig. 3c) (see also Gómez-Escolar et al., 2014; Tao et al., 2015).

In comparison with the SSWs, the variations of $\bar{\omega}^*$ during VI onset (Fig. 3b) are less well synchronized with that in the Arctic (Fig. 1d), perhaps due to the relative weakness of the wave driving and also due to influences from the QBO. Although $\bar{\omega}^*$ is quite noisy, temperatures during VI onset show significant warming in the tropical lower stratosphere (Fig.

3d), probably related to adiabatic warming from anomalous downwelling (Fig. 3b). By mid-February, a downward
290 propagating cooling anomaly can be seen in the tropical upper stratosphere (Fig. 3d), as one would expect from the
anomalous upwelling (Fig. 3b). As noted before, FWs after VIs are dynamically similar to SSWs, and this is also noticeable
in the Tropics. For example, the enhanced extratropical wave driving at the FW is also reflected in the tropical $\bar{\omega}^*$.

4.2 QBO Influences on Tropical Ozone

Understanding the changes in tropical ozone in response to Arctic stratospheric circulation events is complicated by the
295 simultaneous influences from the QBO. To disentangle the two effects, we first examine how the vertical structure of
tropical ozone changes in response to the QBO. Fig. 4a shows the vertical cross-section of tropical ozone anomalies ($\pm 15^\circ$)
composited on the phase of the QBO from 16 QBO cycles. The black curve represents the mean evolution of UEQ30, where
a QBO cycle is defined by two consecutive maxima in UEQ30. Assuming a mean QBO period of 28 months (Baldwin et al.,
2001), a one-degree phase change of the QBO corresponds to ~ 2.3 days. Tweedy et al. (2017) performed a similar analysis
300 (their Fig. 1) by defining the central month of a QBO cycle from changes in the vertical wind shear at 40 hPa and taking
QBO composites for different lags. Our results (Fig. 4a) are in good agreement with their study, e.g., there is a nodal point of
small ozone variations between 10 and 20 hPa, with much stronger variations above and below. Our result also agrees with
Baldwin et al. (2001), that maximum column ozone values occur when the westerly wind shear descends into the lowermost
stratosphere. The vertical structure of the QBO ozone anomalies in Fig. 4a also shows two maxima at ~ 10 hPa and ~ 30 hPa,
305 shifted by about a quarter QBO cycle, consistent with previous findings (Coy et al., 2016; Randel and Wu, 1996).

Fig. 4b demonstrates that SSWs and VIs occur during virtually any phase of the QBO, making it difficult to cleanly
separate the ozone changes from the Arctic and the QBO. However, as shown by the mean timing of the events (V and S
markers at the right), there is a slight preference for SSWs to occur during the easterly QBO phase and VIs during the
westerly QBO phase, a possibility that was discussed by Dunkerton et al. (1988). To filter out the QBO influences from the
310 tropical ozone, we define the QBO ozone signal as the mean ozone anomalies over day -60 to day -30 with respect to the
SSW/VI central date, which is then subtracted from the ozone associated with each Arctic circulation event. We use the
resulting ozone anomalies for preparing Figs. 5c and 5d.

4.3 Tropical Ozone Changes

Fig. 5 presents composite anomalies and composite anomalous tendencies in tropical ozone during SSWs and VIs. The
315 variations in tropical column ozone are rather small and amount to only $\sim 0.5 - 1\%$ of the climatological values, which can be
compared to the $10 - 15\%$ changes seen over the Arctic. Nevertheless, the changes in tropical ozone are quite coherent and
persistent. For SSWs, the column ozone tendency (Fig. 5a, green line) is as expected anti-correlated with the extratropical
EP flux (Fig. 1a), suggesting a direct influence of SSWs on tropical ozone through enhanced residual circulation. SSWs are
followed by a small reduction in tropical column ozone by ~ 2.5 DU ($\sim -1\%$) and an increase by $\sim 1-2$ DU ($\sim 0.5\%$) after mid-
320 March, which persists until late spring. Fig. 5c shows the vertically resolved composite for tropical ozone after removing the

preexisting ozone signal from the QBO, indicating that the local tropical ozone anomalies associated with SSWs are confined to levels above ~60 hPa. During SSW onset, the response of ozone is characterized by significant increases in the upper stratosphere and decreases below the middle stratosphere (~10 hPa), roughly opposite to that in the Arctic (Fig. 2c). The ozone anomalies reverse sign after mid-February and persist into late spring.

325 During VIs (Fig. 5b), there are small tropical column ozone anomalies, which are mostly positive (~1 DU or 0.5%) and
only become negative (~2 DU or 1%) after the FW. However, the vertically resolved ozone anomalies with the QBO
influence removed (Fig. 5d) show a weak dipole in the middle stratosphere around the onset, with little response in the lower
stratosphere. This indicates that the increased column ozone anomalies in Fig. 5b are likely due to the QBO. As discussed
before, the weak tropical ozone response to VIs is linked to the relative weakness of the wave driving during VIs, which is
330 not sufficient to affect the tropical upwelling. However, during the FW of VIs, the wave driving anomaly is strong enough;
the resulting tropical ozone response is similar to that during SSW onset, with a strong and persistent dipole centered at ~20
hPa.

The dynamical mechanisms that create the changes in tropical ozone are dominated by vertical advection associated with
changes to the residual circulation (Randel, 1993). Enhanced tropical upwelling during SSW onset (Fig. 3a) combined with a
335 vertical background of ozone mixing ratios that maximize in the middle stratosphere create positive tendencies above 10 hPa
and negative tendencies below 10 hPa (Fig. 5g). Following the reversal of the residual circulation anomalies at about 10 days
after onset (Fig. 3a), the vertical advection term leads to oppositely-signed ozone anomalies starting at about mid-February.
During VIs, the tropical ozone tendencies (Fig. 5f) are mostly small. There are negative tendencies from vertical advection
(Fig. 5h) in the upper stratosphere and during onset, owing to the weakened meridional circulation from the VI. However,
340 these negative tendencies are compensated by the chemical source term (not shown), overall leading to little changes in
ozone. As expected, the tropical ozone tendencies during the FW of VIs (Fig. 5f) are mostly due to vertical advection (Fig.
5h) and compensating influences from the source term \bar{S} (not shown).

5 Summary and Conclusion

We used MERRA-2 reanalysis to document the composite spatiotemporal ozone response to Arctic circulation events. While
345 the ozone response in the Arctic to Sudden Stratospheric Warming (SSW) events has already been the target of some
previous studies (Butler et al., 2017; de la Cámara et al., 2018b; Hocke et al., 2015), we took a more holistic approach and
studied stratospheric ozone in the Arctic and the Tropics, and we considered not only SSWs but also Vortex Intensification
(VI) and Final Warming (FW) events.

In the Arctic, the onset of SSWs leads to a rapid increase of total ozone by ~50 DU, which over the course of ~60 days
350 gradually transitions towards climatology before the subsequent FWs. Diagnostic analysis using the TEM tracer transport
equation indicates that through the entire life cycle of SSWs, ozone transports by eddies prevail over vertical transports from
the anomalous mean meridional circulation. In contrast, during VIs, Arctic ozone exhibits a slow but progressive decrease,

355 which begins in early winter and results in a ~40 DU reduction by mid-March. The strongest negative ozone tendencies take place right after the central date of VIs, attributable to weakened vertical transports in the lower stratosphere and decreased eddy transports in the upper stratosphere. VIs conclude the winter with a relatively early and strong FW, resembling a mid-winter SSW in terms of the dynamics and ozone perturbations. In contrast, FWs that follow SSWs are relatively late and less remarkable, representing mostly a smooth transition according to climatology. SSWs have also distinct ozone impacts in the Tropics. By removing signals attributable to the QBO, we found tropical ozone responses to SSWs that are largely concurrent and inverse to their Arctic counterparts. At SSW onset, tropical ozone decreases below 10 hPa and increases above, with an opposite behavior after ~20 days when the residual circulation reverses and persists toward the FW. VIs show some obscure tropical ozone responses during onset, presumably due to the relatively weak planetary wave driving anomalies. However, during the FW, VIs are associated with pronounced tropical ozone anomalies due to enhanced vertical transports.

365 There are also some limitations to this study. In terms of the mechanisms, we were mostly focused on the various dynamical effects in changing ozone. However, chemical effects are likely to play also some role in perturbing ozone, in particular in the chemically-dominated upper stratosphere. We were unable to investigate the chemical effects because of the large uncertainties associated with the chemical term in the MERRA-2 reanalysis, but we suspect that the dynamics are overall more important than the chemistry. This is supported by Isaksen et al. (2012), who found that the chemical effect explained only 23% of the Arctic ozone loss during the VI from 2011. Nevertheless, it would be interesting to evaluate the relative contributions from the dynamics and the chemistry in changing ozone during SSWs and VIs, using output from a range of coupled chemistry climate models (CCMs), similar in spirit to de la Cámara (2018b) for SSWs using the WACCM model. We also did not explicitly consider so-called Downward planetary Wave Coupling events (DWCs) (Lubis et al., 2017), relatively short-lived events (< 10 days) associated with increases in ozone before and decreases during the event, leading to a relatively small net response. Our VI events also need to be distinguished from so-called reflective winters, introduced by Shaw and Perlwitz (2013) and discussed by Lubis et al. (2017) to indicate winters in which wave reflection dominates. Although defined in different ways, there is some overlap between years with VIs and reflective winters and they are both associated with negative anomalies in wave driving and ozone.

375 One of the novel results of this study is that FWs that follow VIs induce a surprisingly strong ozone response, which resembles in many respects that of mid-winter SSWs. Another relatively new aspect of this study is that Arctic circulation events also perturb ozone in the Tropics, which is most pronounced during SSWs and early FWs after VIs. This adds to an increasing body of evidence that the mean meridional circulation communicates the effects of Arctic stratospheric circulation events into the lower latitudes. This leads to the notion that the Arctic circulation extremes have an almost global reach, as also evidenced by their impacts on equatorial stratospheric temperatures (Dhaka et al., 2015) and tropospheric equatorial convective activity (Kodera, 2006). It still remains to be seen how the tropical circulation is affected by the combined heating effects from the tropical ozone and the meridional circulation.

Recent studies have suggested that the dynamical coupling between the stratosphere and the troposphere and the surface impact of this coupling is simulated more strongly in models with interactive ozone chemistry (i.e., CCMs) (Haase and Matthes, 2019; Li et al., 2016; Romanowsky et al., 2019), suggesting that intraseasonal variations of ozone are important for the prediction of short-term climate. The results from our study could serve as a reference for the validation of CCMs. Simulations with CCMs in turn could be used to clarify some of the still open questions of the present study, in particular about the response of tropical ozone during VIs and the relative role of photochemistry in changing ozone during the circulation events.

Data availability

MERRA-2 reanalysis are available online via NASA's Goddard Earth Sciences Data and Information Services Center archive (https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/data_access/, Bosilovich et al., 2015).

Author contribution

HH performed the analysis and wrote the manuscript. TR designed the study, provided guidance in the interpretation of the results, and reviewed the manuscript.

Competing interests

The authors declare that they have no conflict of interest.

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Table 1. Central dates t_0 of SSWs and VIs. Numbers in parentheses indicate period (in days) between the central date and the following FW, i.e., $t_{FW}-t_0$.

No.	SSW Central Date	VI Central Date
1	24 Feb 1984 (61)	3 Jan 1983 (88)
2	1 Jan 1985 (82)	19 Jan 1993 (83)
3	23 Jan 1987 (99)	17 Feb 1994 (44)
4	21 Feb 1989 (54)	29 Jan 1996 (72)
5	26 Feb 1999 (66)	29 Jan 1997 (91)
6	11 Feb 2001 (88)	9 Jan 2005 (62)
7	17 Feb 2002 (77)	7 Feb 2011 (57)
8	18 Jan 2003 (86)	6 Jan 2016 (59)
9	5 Jan 2004 (116)	
10	21 Jan 2006 (106)	
11	24 Feb 2007 (54)	
12	22 Feb 2008 (69)	
13	24 Jan 2009 (106)	
14	6 Jan 2013 (117)	
15	12 Feb 2018 (63)	
Mean	3 Feb (83)	23 Jan (70)

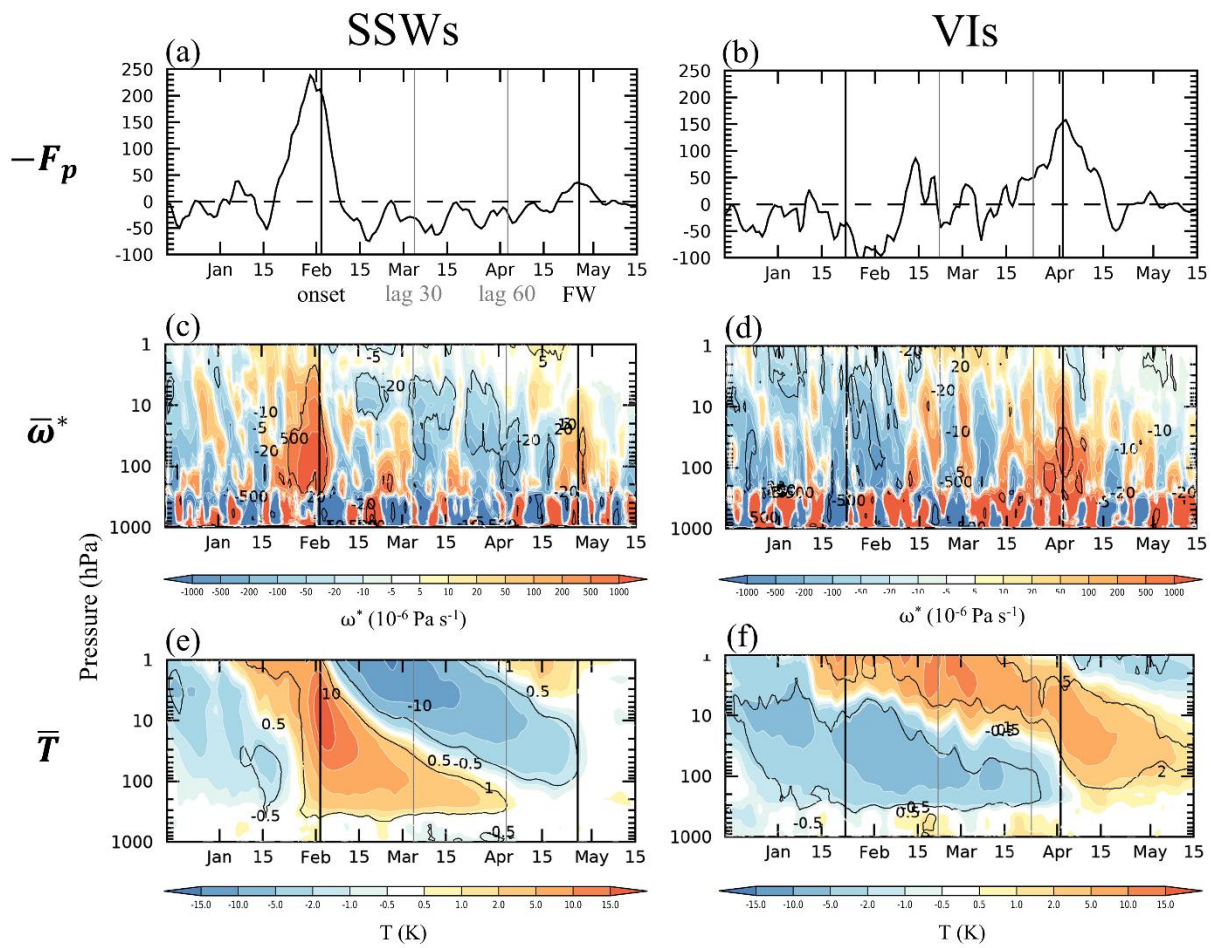


Figure 1. SSW (left) and VI (right) composites over the Arctic. Shown are (a-b) time series of 10-day smoothed vertical EP flux (10^4 kg m s^{-4}) averaged over 40° N - 80° N at 100 hPa, and time-height cross-sections for (c-d) vertical component of the residual circulation ($10^{-6} \text{ Pa s}^{-1}$) (65° N - 85° N) and (e-f) temperature (K) (65° N - 90° N). Contours represent statistical significance at the 95% level.

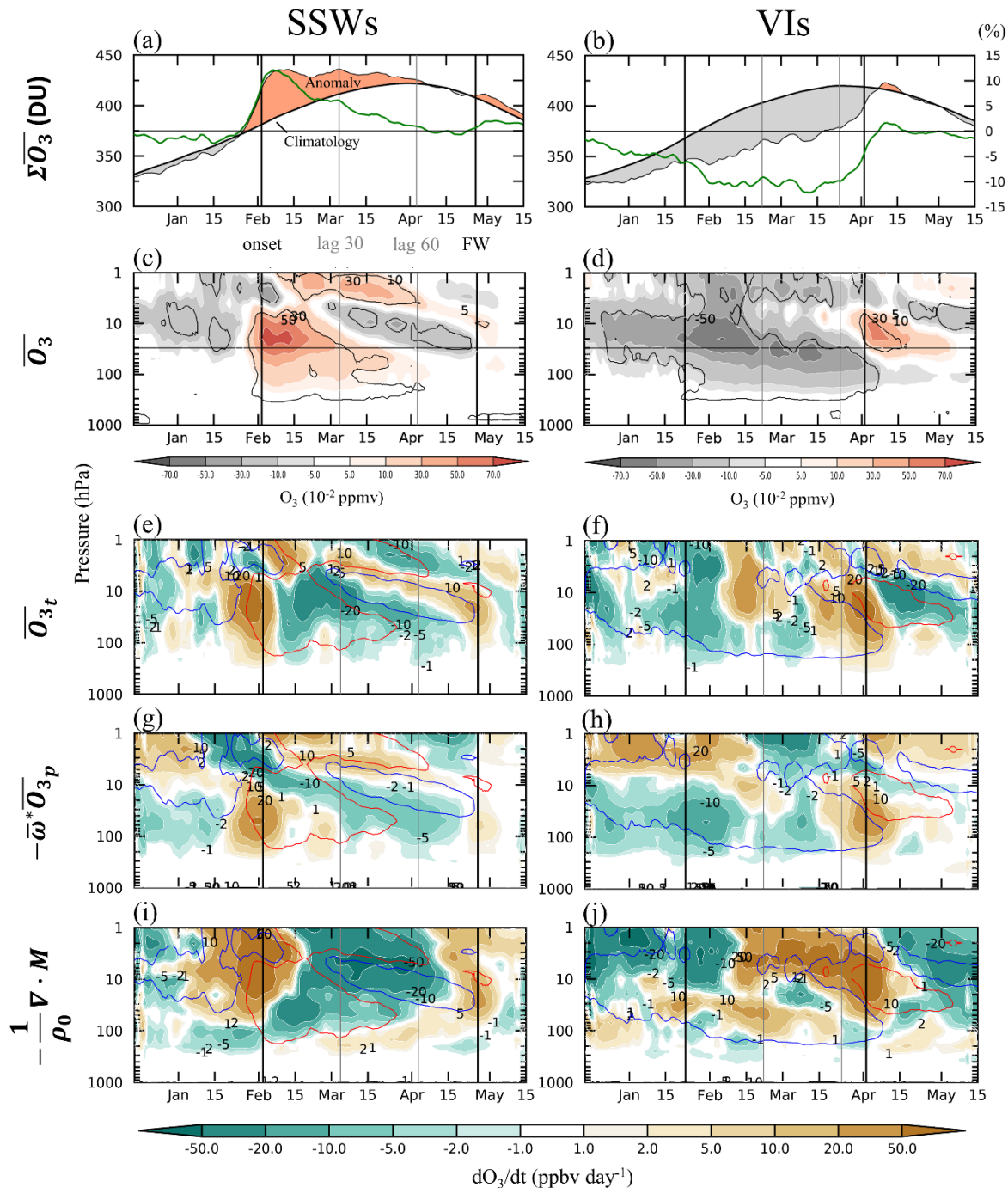


Figure 2. Arctic ozone composites during (left) SSWs and (right) VIs. (a-b) Column ozone (left axis) and associated **percent anomalies with respect to climatology** (right axis); the horizontal line is zero **anomaly**. Remaining panels are **anomalous** time-height cross-sections of (c-d) ozone mixing ratio (10^{-2} ppmv), (e-f) overall ozone tendency, ozone tendency due to (g-h) vertical advection and (i-j) eddy flux convergence (ppbv day^{-1}). Quantities are averaged over 65°N - 90°N for ozone and 65°N - 85°N for

tendencies. Horizontal lines in (c-d) mark the 30 hPa level and contours represent statistical significance at the 95% level. Contours in (e-j) represent the ± 0.1 ppmv ozone anomalies from (c-d).

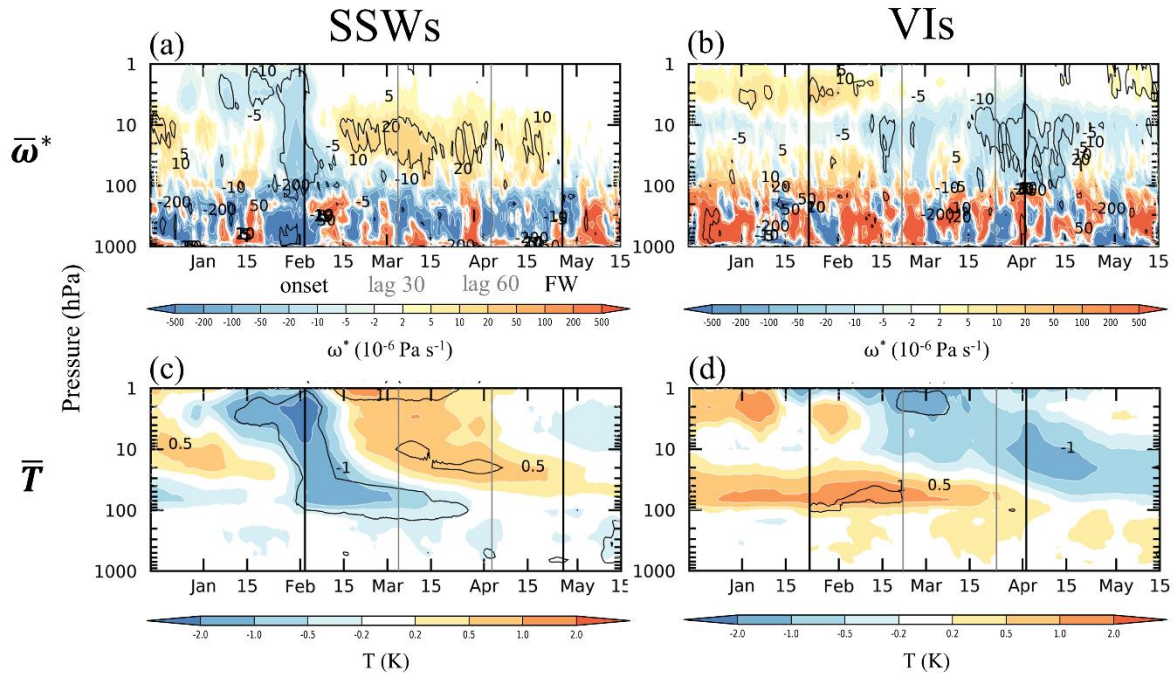


Figure 3. Composite anomalies for (left) SSWs and (right) VIs over the tropical belt ($\pm 15^\circ$). Shown are time-height cross-sections for (a-b) the vertical component of the residual mean circulation ($10^{-6} \text{ Pa s}^{-1}$) and (c-d) temperature (K). Contours are as in Fig. 1.

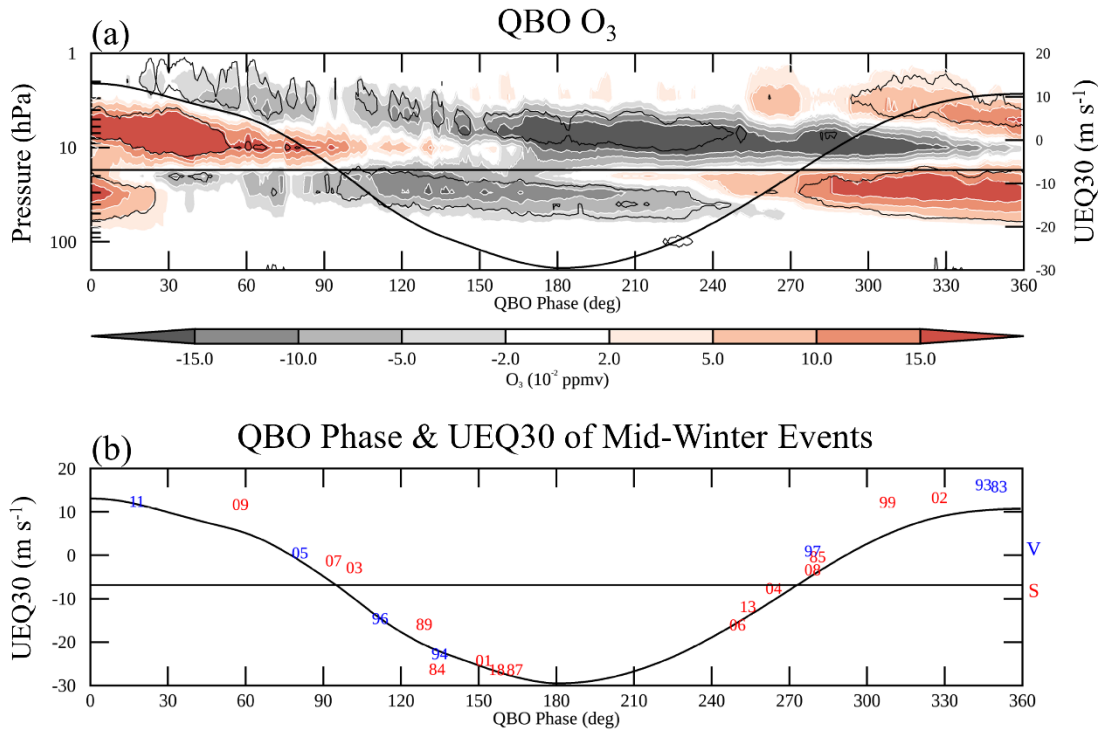


Figure 4. Composites for QBO events. (a) QBO influences on tropical ozone; shading shows composite tropical ozone anomalies ($\pm 15^\circ$) from 16 QBO cycles (1980-2018); black contours represent statistical significance at the 95% level. A QBO cycle is defined by two consecutive positive UEQ30 maxima. (b) Central date timing of selected mid-winter stratospheric circulation events relative to the QBO phase. Red (blue) numbers indicate years and QBO phase of SSWs (VIs); S and V on the right axis is the mean UEQ30 of all SSWs and all VIs (except 2016), respectively. The 2015-2016 QBO event has been purposefully excluded from this analysis due to the anomalous nature of this event. Horizontal line is the climatological mean UEQ30.

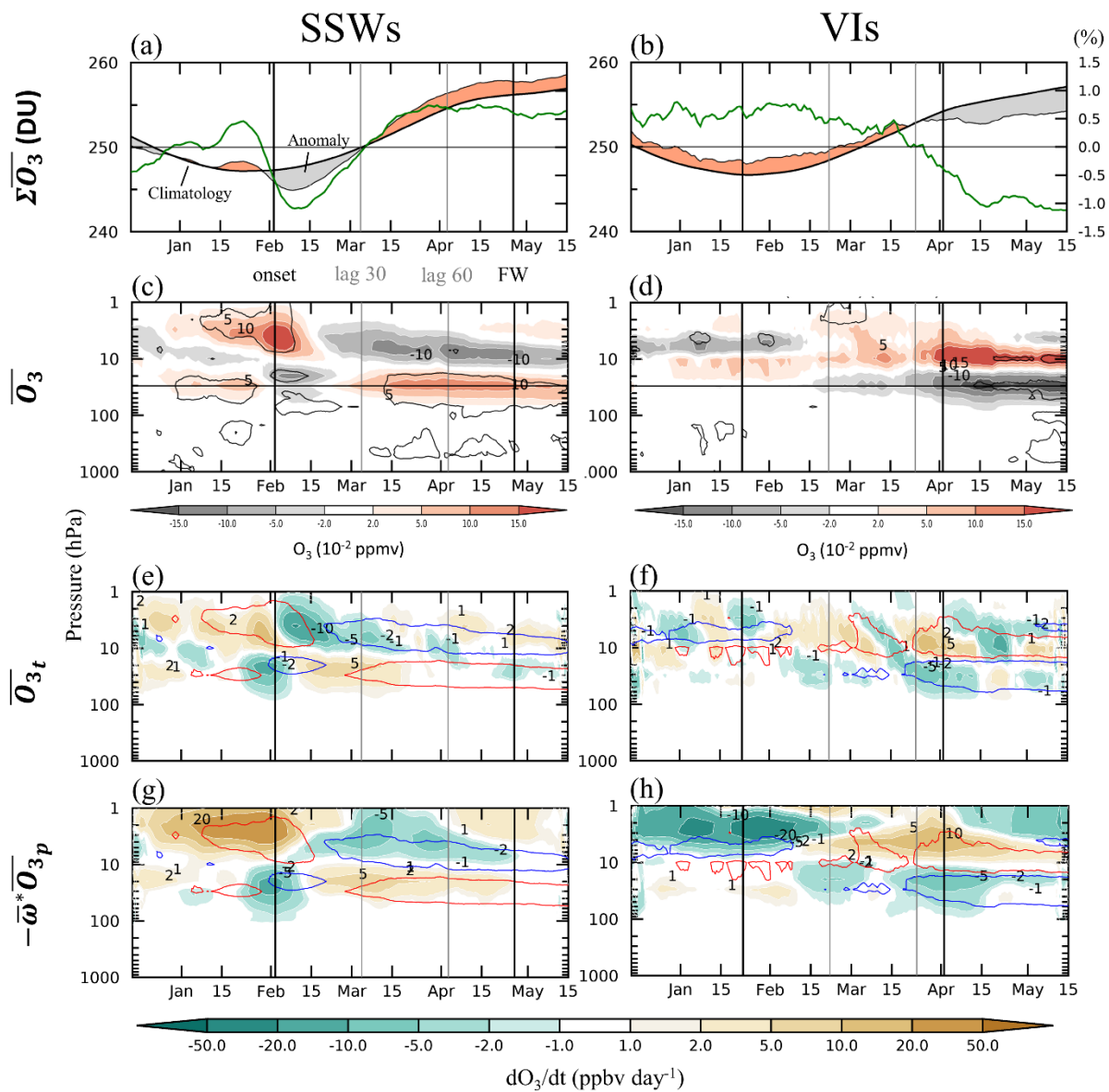


Figure 5. As Fig. 2, except for tropical ozone ($\pm 15^\circ$) and the exclusion of the eddy flux convergence term. Contours in (e-h) are the ± 0.05 ppmv ozone anomalies from (c-d).