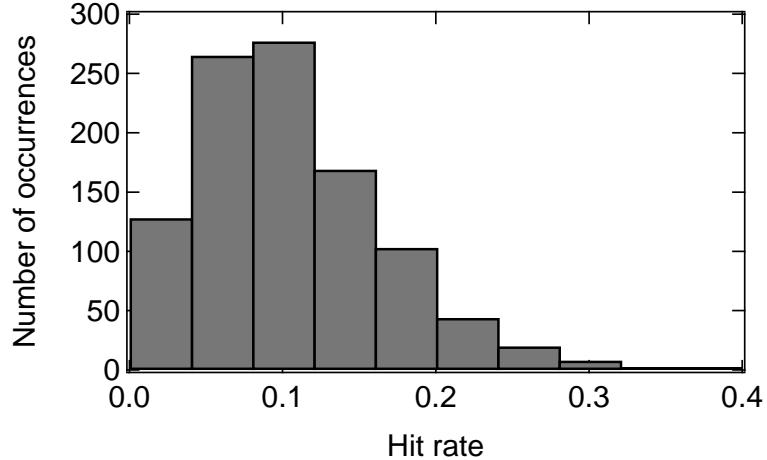


**Figure S1.** Satellite image (visible range from MODIS) from 4 July 2014 with a compilation of flight tracks conducted during 4 - 21 July 2014 (indicated with different colors). On July 21, two flights were performed indicated by a and b.

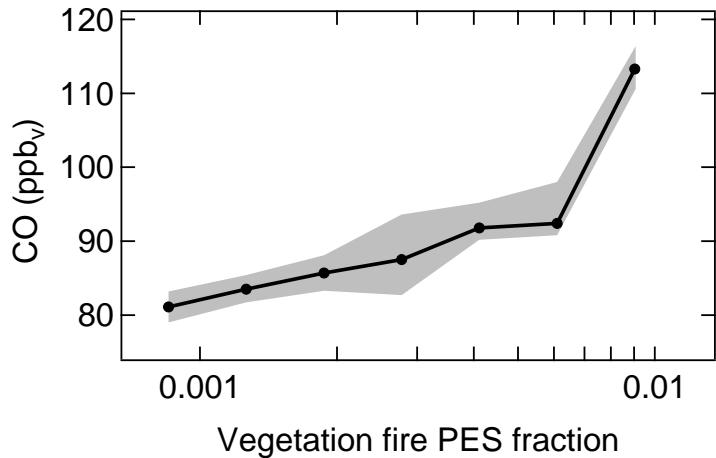
## 1 Hit rate variability during NETCARE 2014

The following approach estimates the NETCARE 2014 hit rate variability that is controlled by several instrumental issues. Figure S2 presents the distribution of the averaged hit rate measured during the Arctic field experiment outside clouds. By



**Figure S2.** Histogram of hit rate values with each value includes  $N = 100$  trials (here, laser shots), measured during the Arctic field experiment (NETCARE 2014) outside clouds. The averaged hit rate value is 0.1 ( $p = 0.1$ ).

assuming that the measured distribution cannot solely be explained by binomial distributed numbers, it needs to be described by a combination of both binomial distributed numbers and the variability caused by instrumental issues. Combined with the



**Figure S3.** CO mixing ratios as a function of FLEXPART predicted air mass residence time over vegetation fires during the southern air mass period. The graph presents the median (solid line with marker) and interquartile ranges (shaded area) of the CO mixing ratios.

Gaussian propagation of uncertainties, the relative standard deviation of the measured hit rate distribution ( $\sigma_{\text{measured}}$ ) can thus be estimated as follows:

$$\sigma_{\text{measured}}^2 = \sigma_{\text{binomial}}^2 + \sigma_{\text{others}}^2, \quad (1)$$

with  $\sigma_{\text{measured}} = 0.06$ , the relative standard deviation related to instrumental issues ( $\sigma_{\text{others}}$ ), and the relative standard deviation of the Binomial probability distribution ( $\sigma_{\text{binomial}}$ ), which is calculated as follows:

$$\sigma_{\text{binomial}} = \frac{\sqrt{N \cdot p \cdot q}}{N}, \quad (2)$$

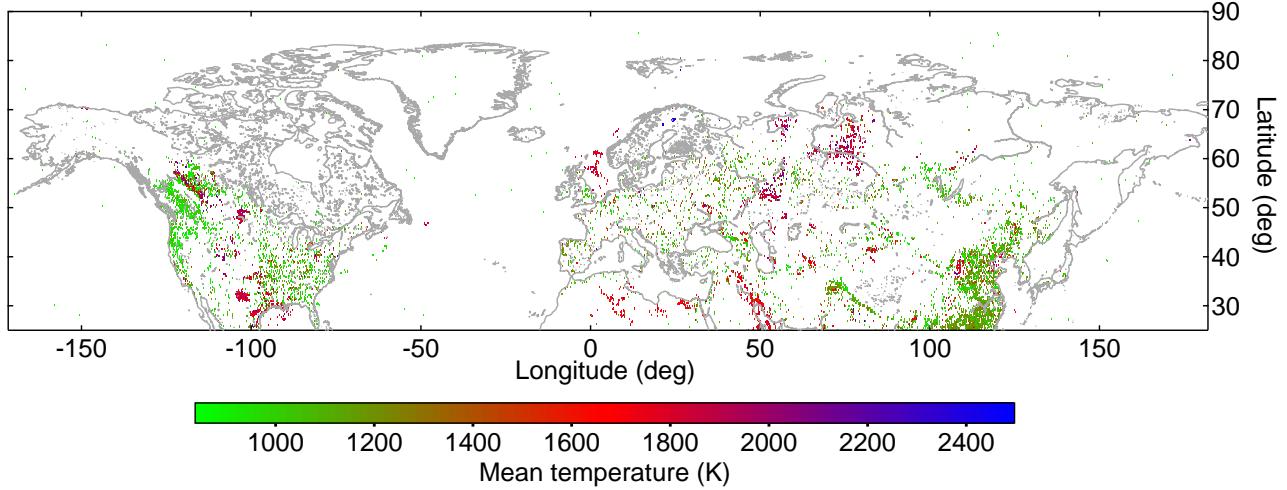
with trials ( $N$ ) is 100, probability of success ( $p$ ) is 0.1, and probability of failure ( $q$ ) is 0.9. The resulting  $\sigma_{\text{binomial}}$  is 0.03. Based on  $\sigma_{\text{binomial}}$  and Eq. 1, the  $\sigma_{\text{others}}$  is 0.05. Together, since the averaged ALABAMA hit rate of 0.1 can vary by 0.05 due to instrumental issues, it is assumed for the analyses in Sects. 3.3 - 3.5 that the particle fraction needs to vary by more than a factor of 2 to be significant.

## 2 Source sectors

### 2.1 Vegetation fires

The Visible Infrared Imaging Radiometer Suite (VIIRS) onboard the Suomi National Polar-Orbiting Partnership spacecraft can detect sub-pixel thermal anomalies by signals in mid- and thermal infrared spectral range (Schroeder et al., 2014; Schroeder and Giglio, 2018). Thermal anomalies typically include both active vegetation fires and industrial heat sources by fossil-fuel combustion etc. Data of the so-called near real-time VIIRS 375 m Active Fire product were downloaded online for 19 June until 21 July 2014 (NASA, 2018) and projected on the gridded maps by adding up observations in each grid cell. Thermal anomalies were identified by one or more observations per grid cell within the respective time.

The presence of vegetation fires can be distinguished from other thermal anomalies by the following approach. Anthropogenic heat sources, such as industry, flaring, and densely populated regions, are separately mapped (see following sections) and subsequently excluded from the thermal anomaly map. Residual heat sources in Fig. 4a thus present vegetation fires, accumulated for 19 June until 21 July 2014. This approach looks reasonable in comparison with Arndt et al. (2015), at least for boreal fires



**Figure S4.** Map showing the mean infrared emission temperatures of industrial heat sources for November 2013 until January 2014.

in northern Canada and Siberia. North American and north Asian biomass burning budgets were anomalously high in 2014 compared to 2001 - 2013 (Fig. 2.57 and Table 2.8 in Arndt et al., 2015). Particularly in summer 2014, large fires in the Great Slave Lake region and Siberia were evident (Fig. 2.40 in Arndt et al., 2015), confirming results in Fig. 3a. Consistent with this, Fig. S3 shows a positive correlation between CO mixing ratios and residence time spent over vegetation fires. However, the main weakness of this approach is given by temporal variabilities of the data sets.

Vegetation fire injection heights are highly variable owing to a combination of fire type, fuel consumption, and prevailing meteorological conditions (e.g., Lavoué et al., 2000; Andreae and Merlet, 2001; Leung et al., 2007; Kondo et al., 2011; Val Martin et al., 2018). Several studies show injection heights for boreal fires in North America and Siberia primary between 1 km and 5 km (e.g., Lavoué et al., 2000; Stohl, 2006; Turquety et al., 2007; Leung et al., 2007; Val Martin et al., 2018). In view of having no better information, the footprint layer for vegetation fires is here defined between 1 km and 5 km.

## 2.2 Industrial sources and flaring

The Nightfire product from VIIRS offers quantitative estimates of infrared source emission temperatures by taking advantage of absent solar reflectance in short-wave infrared spectral range during night (Elvidge et al., 2013, 2016). Emission temperatures typically exhibit bimodal distributions by a combination of gas flares dominating the upper mode (peaking in the 1600 - 2000 K range) and other industrial heat sources as well as vegetation fires dominating the lower mode (peaking in the 800 - 1200 K range; Elvidge et al., 2016; Liu et al., 2018). Thus, temperature distributions of vegetation fires and industrial sources can overlap, depending on the industry sector and the vegetation type. However, vegetation fires at mid- to northern latitudes exhibit a clear seasonality with minimum occurrence between November and January (e.g., Giglio et al., 2013). Consequently, this analysis took advantage of both emission temperatures and seasonality to differentiate between the presence of industrial hot spots and vegetation fires (Elvidge et al., 2013; Liu et al., 2018). VIIRS Nightfire data were thus downloaded for November 2013 until January 2014 and for latitudes north of 25°N (NOAA, 2018a). Data points were projected on the gridded map by averaging temperatures in each grid cell (Fig. S4). Industrial hot spots were identified by both pixel mean temperature larger than 840 K<sup>1</sup> as well as one or more observations per grid cell within the respective time, owing to source persistence (Fig. 4b). Gas flares can be clearly detected by VIIRS Nightfire emission temperatures between 1600 K and 2000 K (Fig. S4) independent on season (Elvidge et al., 2016). Based on this temperature definition, VIIRS Nightfire provide the so-called *Flares only* product

<sup>1</sup>840 K presents the lowest mean temperature for industrial sites, indicative for nonmetal-mineral producing plants (Liu et al., 2018).

(NOAA, 2018b). Data between 30 March and 13 May 2014<sup>2</sup>, were projected on the gridded map by adding up observations in each grid cell. Flares were identified by more than one observation per grid cell within the respective time, owing to source persistence (Fig. 4c).

Both approaches look reasonable by comparing Figs. 3b and c with Fig. 8 in Elvidge et al. (2016) and Fig. 8 in Liu et al. (2018). Oil/gas- and coal-related plants in northwestern Siberia (near Norilsk and Novy Urengoy), northern Europe (North Sea, Norwegian Sea, and Nenets region) and Alaska (Prudhoe Bay) can be clearly identified (Figs. 3b and c). However, uncertainties remain by non-continuously operating industrial sources.

### 2.3 Population

Populated areas provide anthropogenic emissions from residential heating, transportation, cooking etc. The criteria for the grid cell to be classified as populated is given by more than 1000 people live in a  $0.25^\circ \times 0.25^\circ$  area. Geolocation data on population are provided online (MaxMind, 2018). Densely populated regions in central Europe and coastal North American can be clearly identified, whereas Northern Canada, Alaska, and Siberia present sparsely populated regions (Fig. 4d).

### 2.4 Arctic natural marine sources

Sea ice coverage in July 2014 is assessed to obtain areas of open water in the Arctic. The Special Sensor Microwave Imager/Sounder (SSMIS) onboard the Defense Meteorological Satellite Program (DMSP) provides data on sea ice coverage, based on brightness temperature (Cavalieri et al., 1996, updated yearly). Data represent the fractional amount of sea ice covering a  $25 \times 25$  km grid cell in a polar stereographic projection. Polar stereographic coordinates were converted to geodetic latitude and longitude for polar regions (Snyder, 1982)<sup>3</sup>. The obtained data grid was synchronized with the gridded map by averaging concentrations in each grid cell. Arctic open water is defined by both sea ice concentrations less than 30 % and latitudes north of  $73.5^\circ$  N (Fig. 3e).

## 3 Representativeness of the July 2014 measurement periods

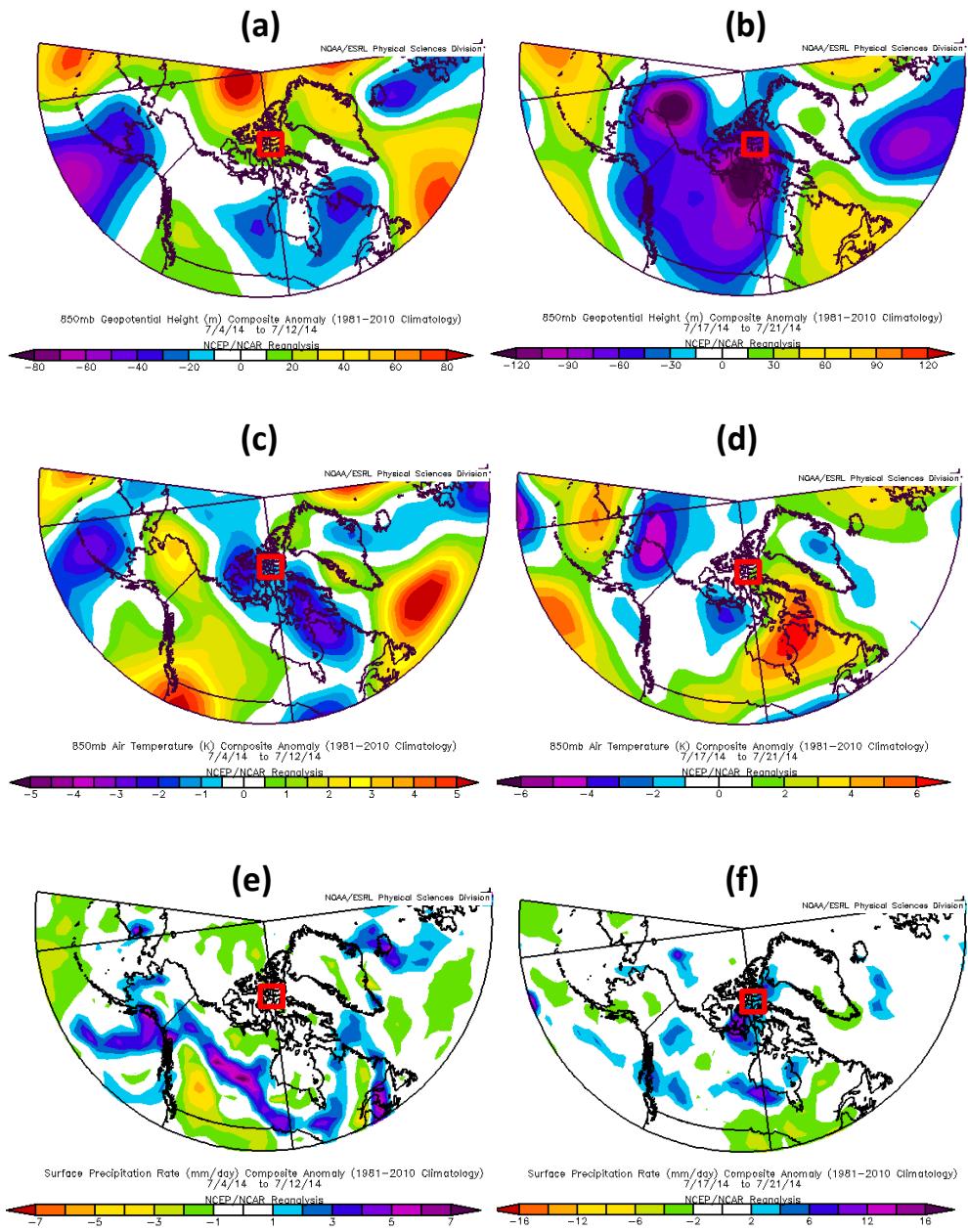
It is of relevance to assess the representativeness of both NETCARE 2014 periods with respect to climatological conditions, i.e. mean conditions between 1981 - 2010. Climatic anomalies were obtained by using re-analysis data prepared by NCEP/NCAR (Kalnay et al., 1996). Images in Fig. S5 were provided by the NOAA/ESRL Physical Sciences Division, Boulder Colorado, from their web site at <http://www.esrl.noaa.gov/psd/data/composites/day/> (Kalnay et al., 1996).

The synoptic situation with the high pressure influence during the Arctic air mass period was generally consistent with the climatology. First, 850 hPa geopotential heights were within +20 m of the climatology in the measurement region (Fig. S5a). Second, air temperatures at 850 hPa were maximum 1 K lower than normal (Fig. S5c). Third, the precipitation anomaly was negligible small (Fig. S5e). Finally, local meteorological conditions, such as the shallow stable stratified BL and low wind speeds (see Sect. 3.1.1), represented typical features in the summertime Arctic atmosphere (e.g., Fuelberg et al., 2010; Tjernström et al., 2012).

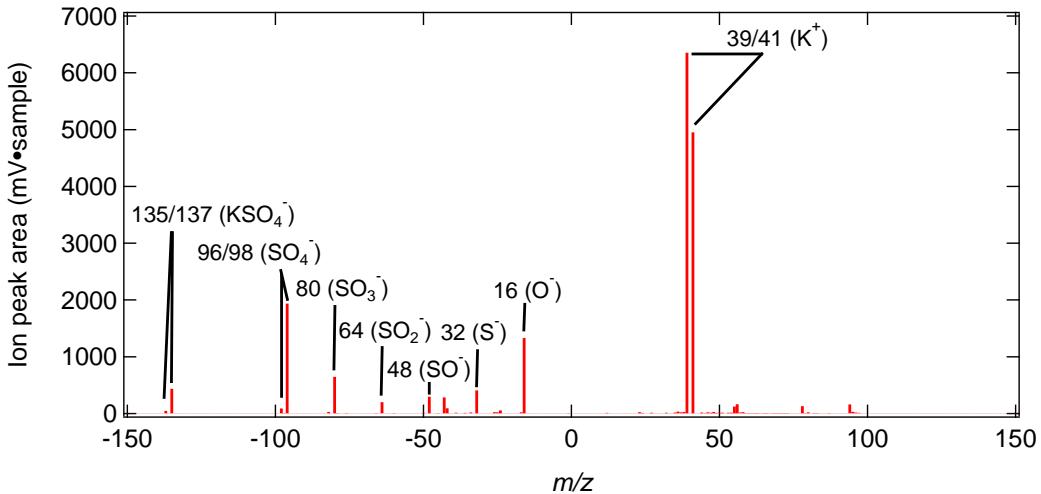
In contrast to this, the southern air mass period with influences of the low-pressure system centered south of Resolute Bay were significantly anomalous compared to climatological mean. Geopotential height at 850 hPa shows a climatological departure down to -120 m close to the center of the low-pressure system (Fig. S5b). Further, warmer air than normal was advected into and more precipitation than normal fell in the measurement region (Figs. S5d and f). Air temperatures at 850 hPa deviated up to +6 K from the climatology in the region north-east of the Hudson Bay (Fig. S5d). Additionally, the vegetation fire activity in North America was anomalously high in summer 2014 compared to 2001 - 2013 (Arndt et al., 2015). Together, influences of vegetation fires on Arctic atmospheric composition were larger than on climatological average caused by a combination of anomalous meteorological conditions and intensive fires in northern Canada.

<sup>2</sup>Data earlier in 2014 are not available and high latitude data later than May are not valuable due to 24 h solar irradiance.

<sup>3</sup>Source code was taken from <ftp://sidads.colorado.edu/pub/DATASETS/seoice/polar-stereo/tools/mapxy.for>.



**Figure S5.** Anomaly from climatological mean (1981 - 2010) of 850 hPa geopotential height (a-b), 850 hPa air temperature (c-d), and 850 hPa precipitation rate (e-f) during the Arctic air mass period (left panel) and the southern air mass period (right panel). The measurement region is marked with a red box.



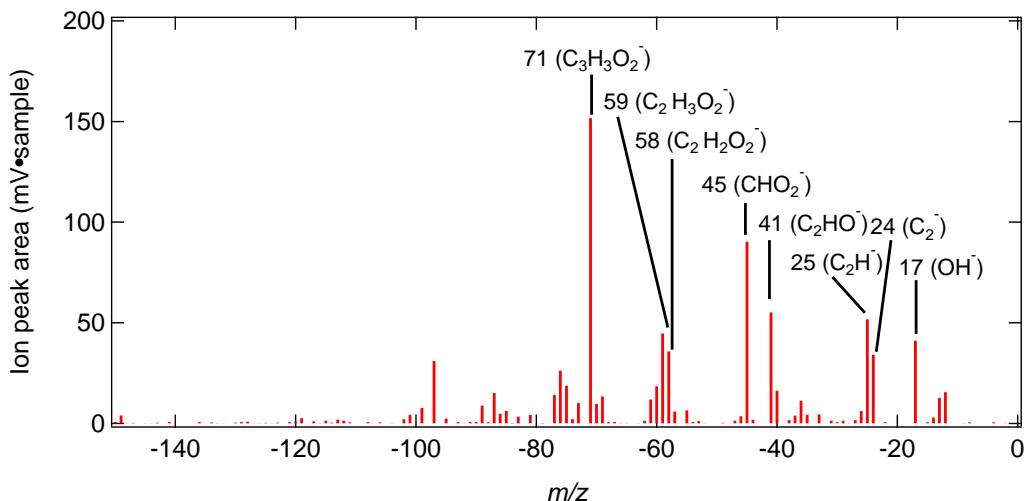
**Figure S6.** Potassium sulfate ( $\text{KSO}_4$ ) mean bipolar spectrum of 69 single particle spectra analyzed by the ALABAMA. The isotopic ratio of potassium at  $m/z +41$  and  $m/z +39$  is too high, likely caused by saturation effects of the MCP.

#### 4 Levoglucosan and Potassium Sulfate

Levoglucosan (or 1,6-anhydro-beta-glucopyranose;  $\text{C}_6\text{H}_{10}\text{O}_5$ ) is a key organic tracer for vegetation fires (Simoneit et al., 1999). However, there is no consensus in SPMS literature about ion markers of levoglucosan. The averaged laboratory mass spectrum by Silva et al. (1999) shows ion peaks at  $m/z$  -45, -59 and -71, while others also used  $m/z$  -73 to identify levoglucosan in ambient mass spectra (Moffet et al., 2008; Corbin et al., 2012). Consequently, ALABAMA laboratory measurements were conducted to identify mass spectra characteristics of levoglucosan.

Levoglucosan is poorly absorbing at 266 nm laser wavelength as found in this laboratory study. Therefore, the better absorbing material  $\text{KSO}_4$  was mixed with levoglucosan. Both substances were suspended in purified water and nebulized with an aerosol generation system. Diffusion dryers filled with silica gel removed water from the particles and the resulting particles were introduced into the ALABAMA. Analogous to levoglucosan mixed with  $\text{KSO}_4$ ,  $\text{KSO}_4$  was separately analyzed to differentiate ion peaks originating from levoglucosan and  $\text{KSO}_4$ .

The mean mass spectrum of  $\text{KSO}_4$  is characterized by molecular ion peaks at  $m/z$  -135/137 (Fig. S6). It further shows potassium peaks at  $m/z$  +39/41 as well as several fragments that contain sulfur and/or oxygen ( $m/z$  -96/98, -80, -64, -48, -21, and -16). In parallel, the mean mass spectrum of levoglucosan is obtained by subtracting the mean spectrum of  $\text{KSO}_4$  from levoglucosan mixed with  $\text{KSO}_4$  (Fig. S7). As a result, levoglucosan is characterized by fragments at  $m/z$  -71, -59, -45, confirming laboratory measurements by Silva et al. (1999). Additional fragments at  $m/z$  -58, -41, and -25 have been also reported by Silva et al. (1999), albeit not dominant in their study. However, these mass spectral characteristics are not unambiguous owing to isobaric interferences with fragments of other oxygen-containing organic compounds (Table S1). It is thus difficult to clearly identify levoglucosan in ambient particle mass spectra.



**Figure S7.** Mean anion spectrum of levoglucosan, obtained as the difference between the mean spectrum of levoglucosan with  $\text{KSO}_4$  (average of 754 single particle spectra analyzed by the ALABAMA) and the pure  $\text{KSO}_4$  mean spectrum (see Fig. S6).

**Table S1.** Organic compounds showing isobaric interferences with fragments of levoglucosan.

Compounds	Ion markers	References
Acetic acid	m/z -59	Lee et al. (2003); Brands et al. (2011)
Methylglyoxal	m/z -71, m/z -59	Zauscher et al. (2013)* Zauscher et al. (2013)*
Acrylate	m/z -71	Zauscher et al. (2013)*
Glyoxylic acid	m/z -73	Lee et al. (2003); Zauscher et al. (2013)
Polycyclic aromatic hydrocarbons	m/z -25	Silva and Prather (2000)
Polycyclic aromatic hydrocarbons	m/z -41	Silva and Prather (2000)
-cont. alkohol functional groups	m/z -41	Silva and Prather (2000)
Formic acid	m/z -45	Lee et al. (2003)
Ethoxide	m/z -45	Lee et al. (2003)
n-propoxide	m/z -59	Lee et al. (2003)
Propanoic acid	m/z -73	Lee et al. (2003)

**Table S1.** \*Zauscher et al. (2013, Supplement) further reported that most oxygen-containing organic standards show small ion peaks at m/z -45 and m/z -59.

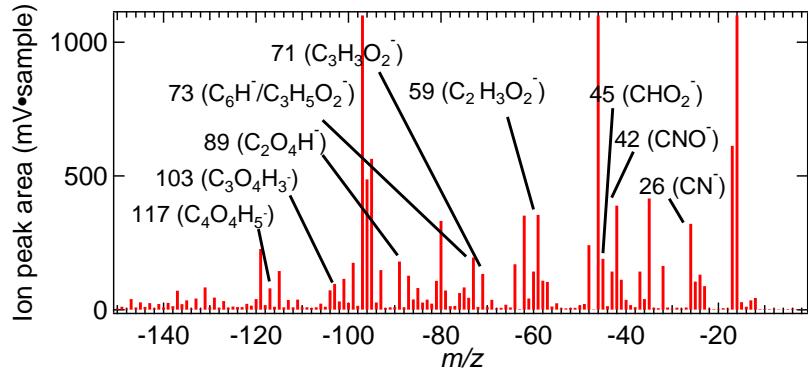
## 5 Internal mixing of particle types

The internal mixing state of the ALABAMA particle types is given in Table S2. Additionally, the expanded mean anion spectrum of those sea spray particles that are internally mixed with DCA compounds are shown in Fig. S8. Figure S9 depicts the bipolar mean spectrum of particles that were not classified in of the particle types introduced in Sect. 3.2.

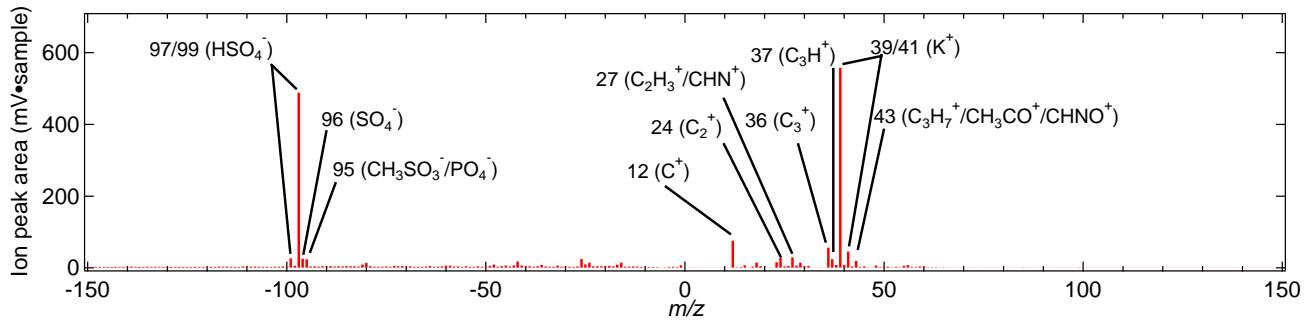
**Table S2.** Particle types and their internal mixing with other substances derived from the mean spectra in Figs. Additional ion signals of sulfate ( $m/z$ -97/99 ( $\text{HSO}_4^-$ ), -96 ( $\text{SO}_4^-$ )) and potassium ( $m/z$ +39/41 ( $\text{K}^+$ )) were present in every mean spectrum and have therefore not been listed here. References from SPMS laboratory and field studies are indicated by numbers that are defined below. These references are used for the assignment of the additional ion signals to the corresponding substances.

Particle type denotation	Ion markers	Additional ion signals in mean spectrum	Corresponding substances
Na/Cl/ Nitrate -containing	$m/z$ +23 ( $\text{Na}^+$ ), +46 ( $\text{Na}_2^+$ ), +62 ( $\text{Na}_2\text{O}^+$ ), +63 ( $\text{Na}_2\text{OH}^+$ ) +81/83 ( $\text{Na}_2\text{Cl}^+$ ) $m/z$ -35/37 ( $\text{Cl}^-$ ), -46 ( $\text{NO}_2^-$ ), -62( $\text{NO}_3^-$ ), -93/95 ( $\text{NaCl}_2^-$ )	$m/z$ +24/25/26 ( $\text{Mg}^+$ ), +40 ( $\text{Ca}^+$ ), +56 ( $\text{CaO}^+$ ) $m/z$ -26 ( $\text{CN}^-$ ), -42 ( $\text{CNO}^-$ ), -45 ( $\text{CHO}_2^-$ ), -59 ( $\text{C}_2\text{H}_3\text{O}_2^-$ ), -71 ( $\text{C}_3\text{H}_3\text{O}_2^-$ ), -73 ( $\text{C}_3\text{H}_5\text{O}_2^-$ ) -89 ( $\text{C}_2\text{HO}_4^-$ )	magnesium <sup>1,2</sup> calcium <sup>1</sup> nitrogen-cont. OM <sup>1,2,3</sup> oxygen-cont. OM <sup>3,4,5</sup> oxalic acid <sup>9</sup>
EC -containing	$m/z$ +12, +24,...,+84 ( $\text{C}_{1-7}^+$ ) $m/z$ -12, -24,...,-96 ( $\text{C}_{1-8}^-$ )	$m/z$ -26 ( $\text{CN}^-$ ), -42 ( $\text{CNO}^-$ ), -45 ( $\text{CHO}_2^-$ ), -59 ( $\text{C}_2\text{H}_3\text{O}_2^-$ ), -73 ( $\text{C}_3\text{H}_5\text{O}_2^-$ ), -89 ( $\text{C}_2\text{HO}_4^-$ ), -95 ( $\text{CH}_3\text{O}_3\text{S}^-$ ), -46 ( $\text{NO}_2^-$ ), -62( $\text{NO}_3^-$ )	nitrogen-cont. OM <sup>6</sup> oxygen-cont. OM <sup>8,10</sup> oxalic acid <sup>9</sup> MSA <sup>9</sup> nitrate <sup>9</sup>
nss-nitrate -containing	$m/z$ -46 ( $\text{NO}_2^-$ ), -62( $\text{NO}_3^-$ )	$m/z$ -26 ( $\text{CN}^-$ ), -42 ( $\text{CNO}^-$ ), -45 ( $\text{CHO}_2^-$ ), -59 ( $\text{C}_2\text{H}_3\text{O}_2^-$ ), -73 ( $\text{C}_3\text{H}_5\text{O}_2^-$ ), -89 ( $\text{C}_2\text{HO}_4^-$ )	nitrogen-cont. OM <sup>6</sup> oxygen-cont. OM <sup>8,10</sup> oxalic acid <sup>9</sup>
TMA -containing	$m/z$ +59 (( $\text{CH}_3)_3\text{N}^+)$ , +58 ( $\text{C}_3\text{H}_8\text{N}^+$ )	$m/z$ +18 ( $\text{NH}_4^+$ ), +12, ..., +36 ( $\text{C}_{1-3}^+$ ), +27 ( $\text{C}_2\text{H}_3^+/\text{CHN}^+$ ), +37 ( $\text{C}_3\text{H}^+$ ), +43 ( $\text{C}_3\text{H}_7^+/\text{CH}_3\text{CO}^+$ / $\text{CHNO}^+$ ) $m/z$ -95 ( $\text{CH}_3\text{O}_3\text{S}^-$ )	ammonium <sup>9</sup> carbon cluster ions <sup>9</sup> hydrocarbons <sup>7</sup> oxidized organics <sup>7</sup> MSA <sup>9</sup>
DCA -containing	$m/z$ -89 ( $\text{C}_2\text{HO}_4^-$ ), -103 ( $\text{C}_3\text{H}_3\text{O}_4^-$ ), -117 ( $\text{C}_4\text{H}_5\text{O}_4^-$ )	$m/z$ +27 ( $\text{C}_2\text{H}_3^+/\text{CHN}^+$ ), $m/z$ -26 ( $\text{CN}^-$ ), -42 ( $\text{CNO}^-$ ), -45 ( $\text{CHO}_2^-$ ), -59 ( $\text{C}_2\text{H}_3\text{O}_2^-$ ), -73 ( $\text{C}_3\text{H}_5\text{O}_2^-$ )	hydrocarbons <sup>7</sup> nitrogen-cont. OM <sup>6</sup> oxygen-cont. OM <sup>8,10</sup>

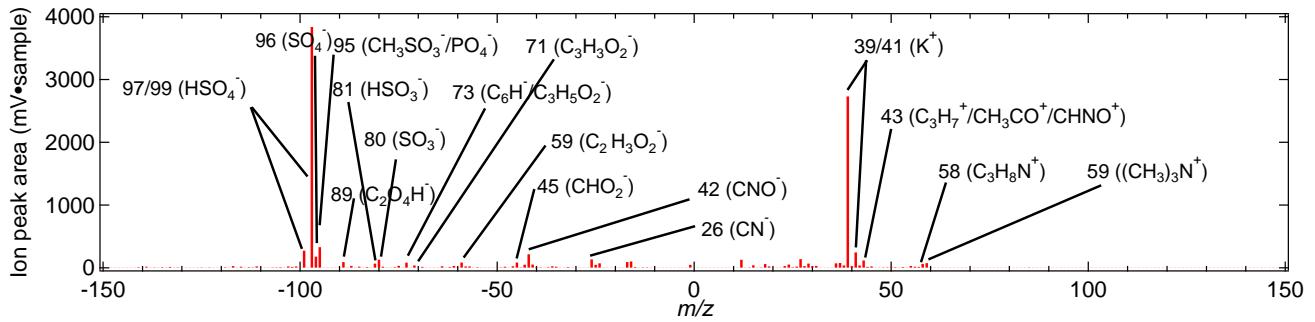
**Table S2.** Given reference numbers are defined as follows: <sup>(1)</sup>Prather et al. (2013),<sup>(2)</sup>Guasco et al. (2014),<sup>(3)</sup>Pratt et al. (2009),<sup>(4)</sup>Schmidt et al. (2017),<sup>(5)</sup>Trimborn et al. (2002),<sup>(6)</sup>Silva et al. (1999),<sup>(7)</sup>Schmidt (2015),<sup>(8)</sup>Zauscher et al. (2013),<sup>(9)</sup>see Table 1, <sup>(10)</sup>Lee et al. (2003).



**Figure S8.** Expanded mean anion spectrum of sea spray particles containing DCA, presenting 60% of all sea spray particles in Fig. 7a.



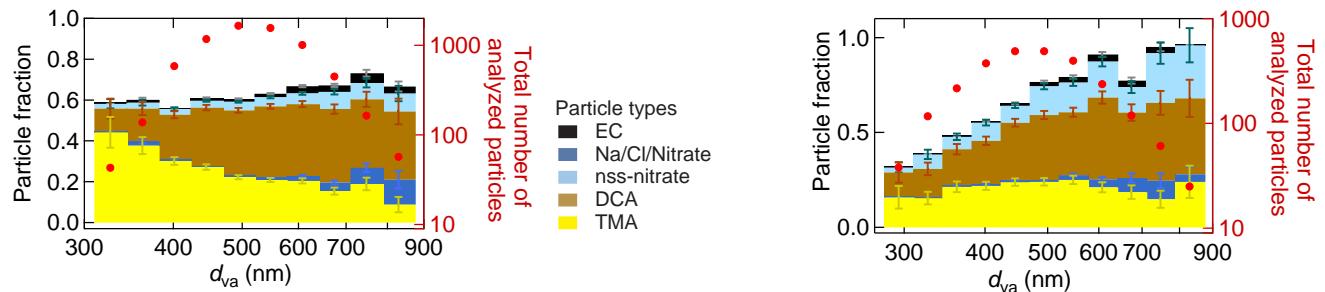
**Figure S9.** Bipolar mean spectrum of the remaining 4689 particles (46 % of all mass spectra analyzed by the ALABAMA) that could not be classified in one of the five distinct particle groups introduced in Sect. 3.2. These particles are therefore summarized as “Others”.



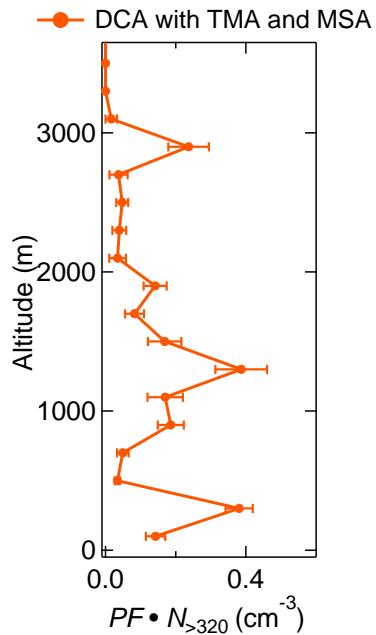
**Figure S10.** Bipolar mean spectrum of particles that show internal mixing of DCA, TMA, and MSA (14 % of all DCA-containing particles, 4 % of all mass spectra analyzed by the ALABAMA).

## 6 Additional figures

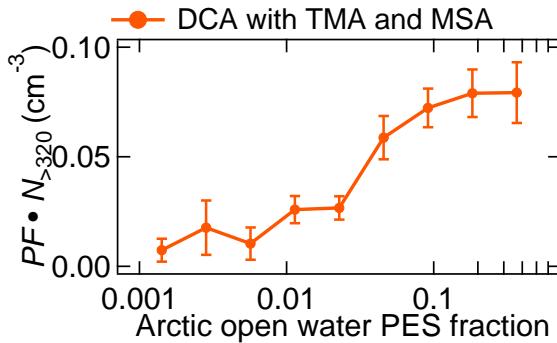
Figures S11 - S15 are associated with the discussion in Sections 3.4 and 3.5 in the main document. Size distributions of the ALABAMA particle types are given in Fig. S11.



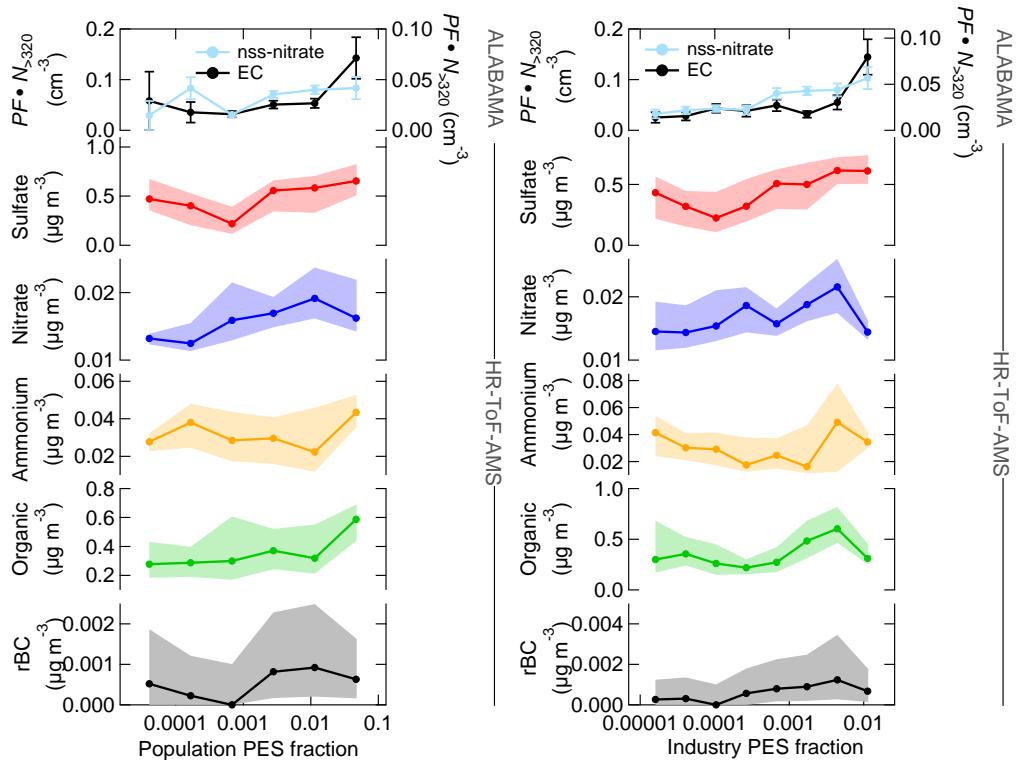
**Figure S11.** Size-resolved aerosol composition of the ALABAMA particle types during the Arctic (a) and southern (b) air mass periods. Both figures represent the cumulative particle fraction of: EC-containing (black), nss-nitrate-containing (light blue), DCA-containing (brown), Na/Cl/Nitrate-containing (dark blue), and TMA-containing (yellow) particles as well as the total number of analyzed particles per bin (red). Uncertainty analysis is given in the Supplement Sect. 7.



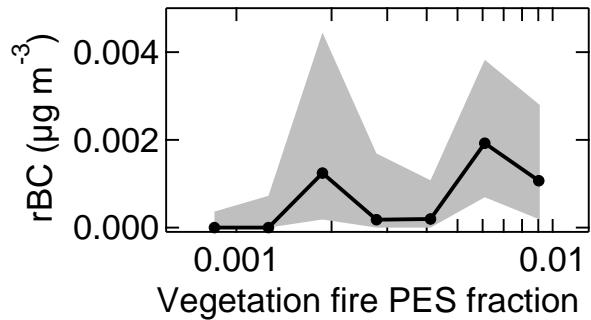
**Figure S12.** Vertically resolved aerosol composition measured by the ALABAMA during the Arctic air mass period. The panel represents the ALABAMA scaled number concentration ( $PF \cdot N_{>320}$ ) of particles that show internal mixing of DCA, TMA, and MSA. Uncertainty analyses are given in the Supplement Sect. 7.



**Figure S13.** Aerosol composition measured by the ALABAMA as a function of FLEXPART derived residence time over Arctic open water areas during the Arctic air mass period. The panel represents the ALABAMA scaled number concentration ( $PF \cdot N_{>320}$ ) of particles that show internal mixing of DCA, TMA, and MSA. Uncertainty analyses are given in the Supplement Sect. 7.

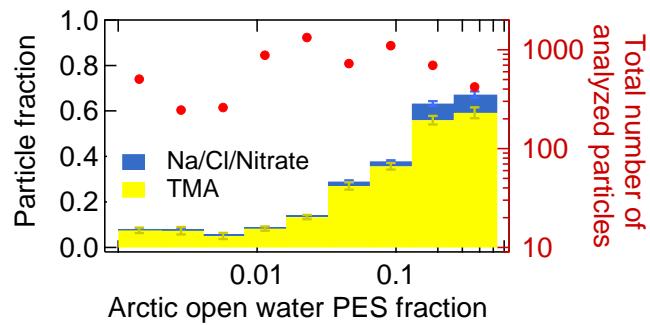


**Figure S14.** Aerosol composition measured by the ALABAMA and by the HR-ToF-AMS (indicated on the right side) as a function of FLEXPART derived residence time over (a) populated areas and (b) industrial areas during the Arctic air mass period. Top panel represents the ALABAMA scaled number concentration ( $PF \cdot N_{>320}$ ) with right axis referring to EC-containing (black) and left axis referring to nss-nitrate-containing (light blue) particles. Lower panels represent each median (solid line with marker) and interquartile ranges (shaded area) of the HR-ToF-AMS mass concentrations of sulfate (red), nitrate (blue), ammonium (orange), and organic matter (light green) as well as rBC measured by the SP2 (black). Uncertainty analyses are given in the Supplement Sect. 7.

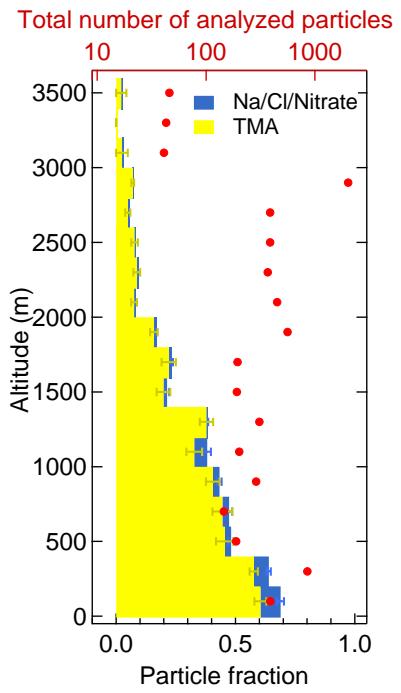


**Figure S15.** Refractory black carbon (rBC) measured by the SP2 as a function of FLEXPART derived residence time over vegetation fires during the southern air mass period. The figure represents median (solid line with marker) and interquartile ranges (shaded area). Uncertainty analyses are given in the Supplement Sect. 7.

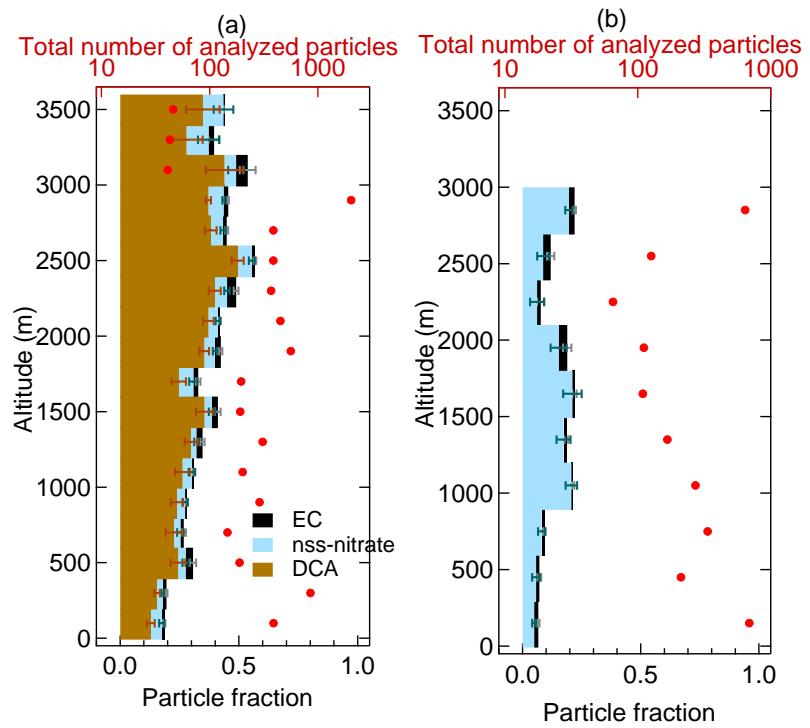
Figures S16-S24 provide the cumulative fraction of the ALABAMA aerosol composition, associated with figures of the scaled number concentration in the main document.



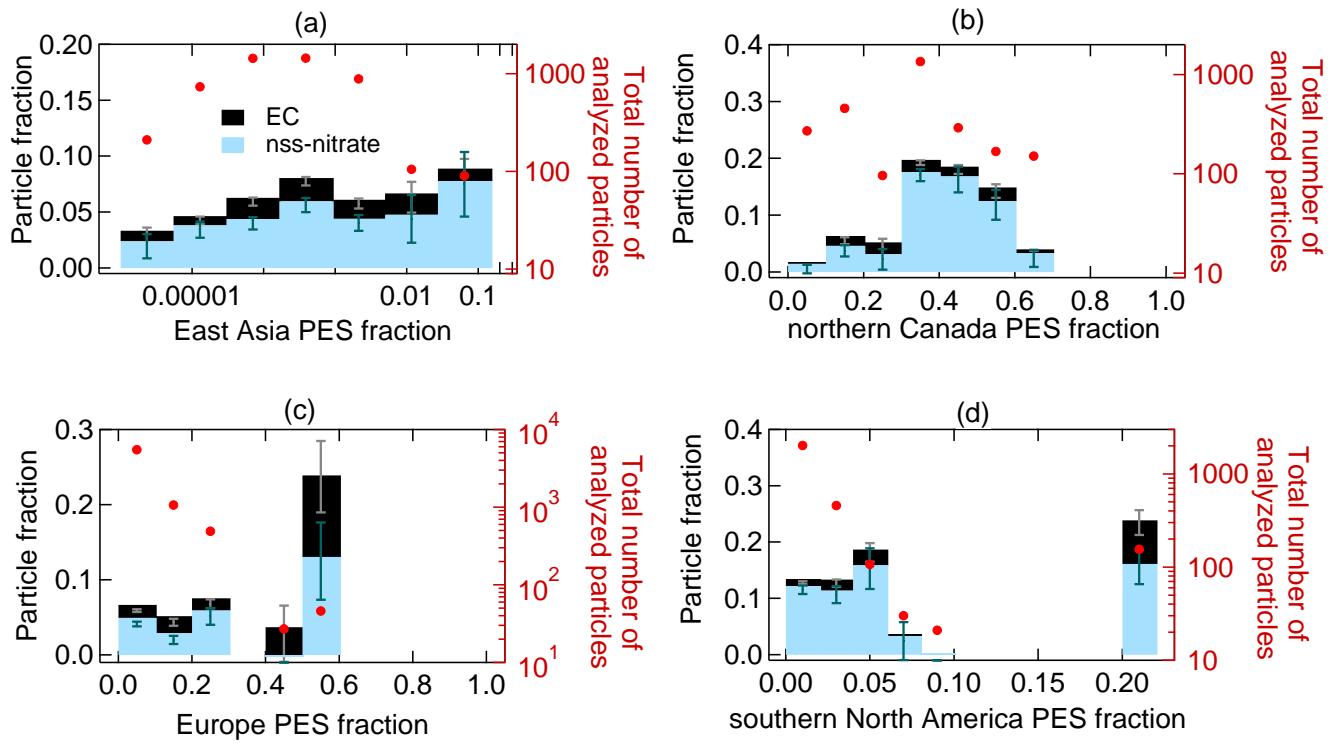
**Figure S16.** Aerosol composition measured by the ALABAMA as a function of FLEXPART derived residence time over Arctic open water areas during the Arctic air mass period. The figure represents the cumulative particle fraction of TMA-containing (yellow) and Na/Cl/Nitrate-containing (blue) particles as well as the total number of analyzed particles per bin (red). Uncertainty analyses are given in the Supplement Sect. 7.



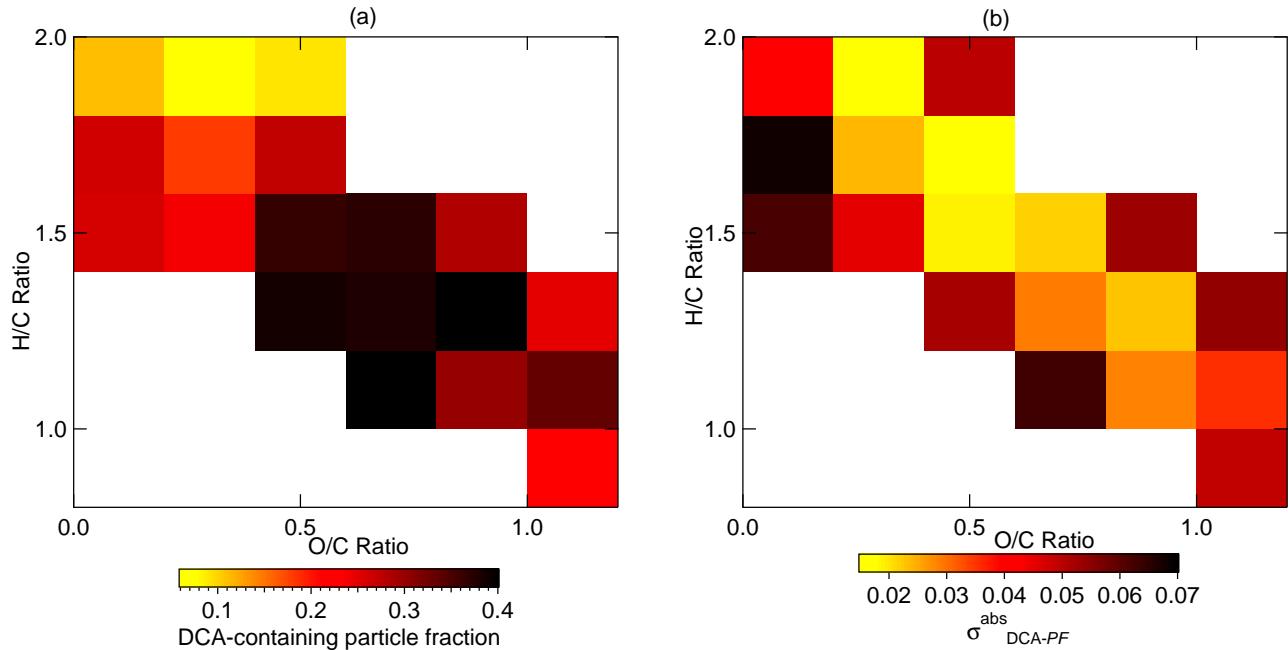
**Figure S17.** Vertically resolved aerosol composition measured by the ALABAMA during the Arctic air mass period. The figure represents the cumulative particle fraction of TMA-containing (yellow) and Na/Cl/Nitrate-containing (blue) particles as well as the total number of analyzed particles per bin (red). Uncertainty analyses are given in the Supplement Sect. 7.



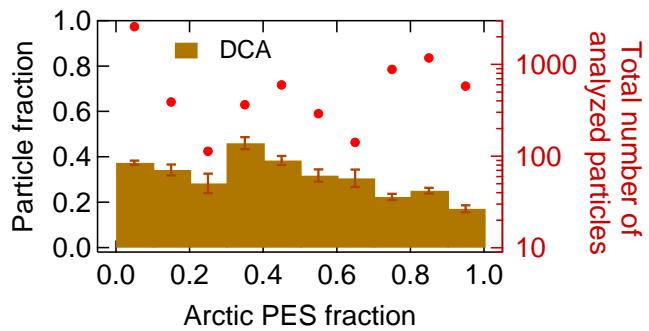
**Figure S18.** Vertically resolved aerosol composition measured by the ALABAMA during (a) the Arctic air and (b) the southern air mass periods. Figures represent the cumulative particle fraction of DCA-containing (brown), nss-nitrate-containing (blue), and EC-containing (black) particles as well as the total number of analyzed particles per bin (red). Uncertainty analyses are given in the Supplement Sect. 7.



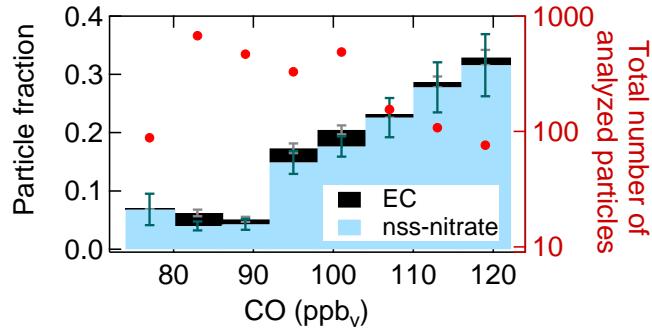
**Figure S19.** Aerosol composition measured by the ALABAMA as a function of FLEXPART derived air mass residence time over the East Asia (a), northern Canada (b), Europe (c), and southern North America (d) during the Arctic (a,c) and the southern air mass (b,d) periods. Figures represent each the cumulative particle fraction of EC-containing (black) and nss-nitrate-containing (light blue) particles as well as the total number of analyzed particles per bin (red). Uncertainty analyses are given in the Supplement Sect. 7.



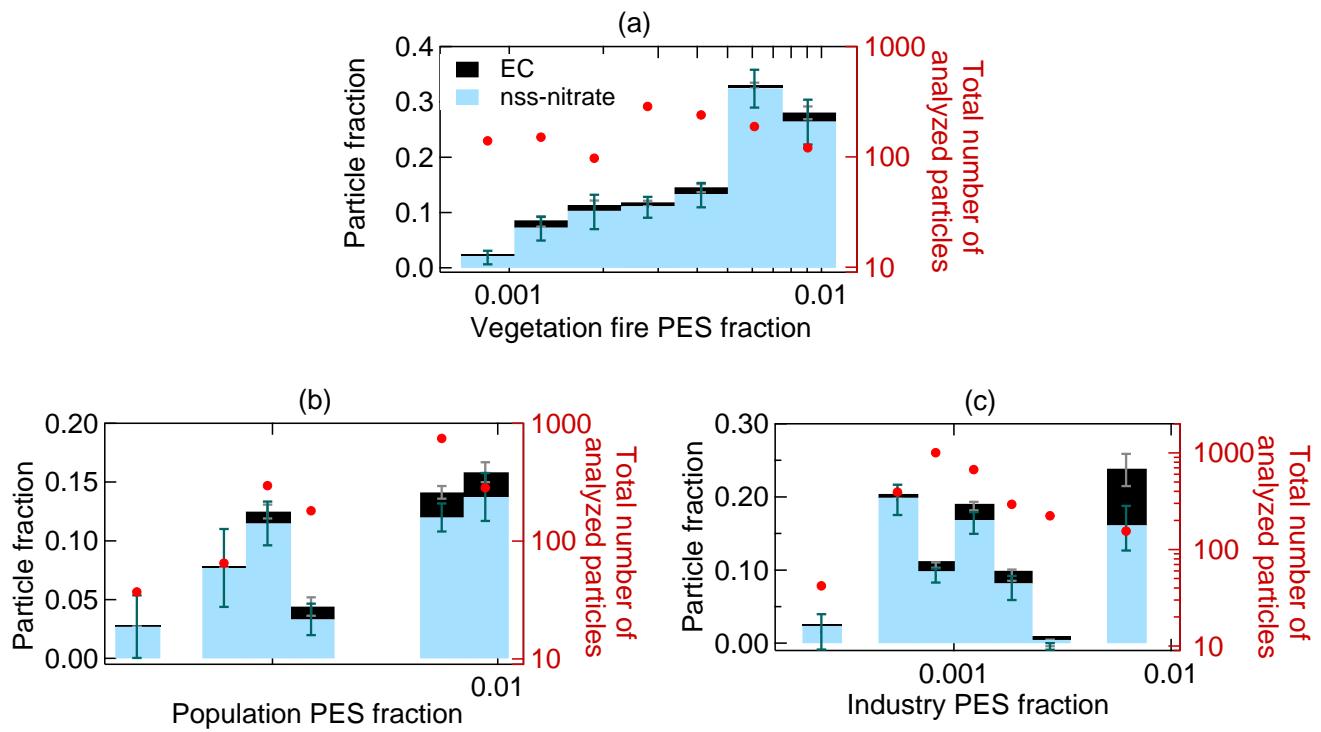
**Figure S20.** Comparison between the HR-ToF-AMS estimated oxygen-to-carbon (O/C) and hydrogen-to-carbon (H/C) ratios colored by (a) the particle fraction of DCA-containing measured by the ALABAMA and (b) the absolute uncertainty of the particle fraction. Uncertainty analyses are given in Sect. 7.



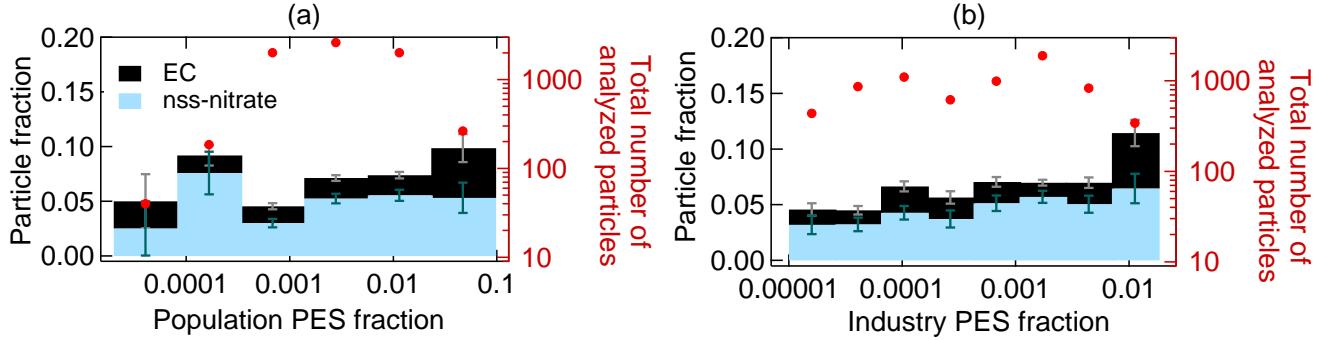
**Figure S21.** Aerosol composition measured by the ALABAMA as a function of FLEXPART derived air mass residence time over Arctic regions during the Arctic air mass period. The figure represents the fraction of DCA-containing particles particles as well as the total number of analyzed particles per bin (red). Uncertainty analysis is given in the Supplement Sect. 7.



**Figure S22.** Aerosol composition measured by the ALABAMA as a function of CO mixing ratios during the southern air mass period. Figure represents the cumulative particle fraction of nss-nitrate-containing (blue) and EC-containing (black) particles as well as the total number of analyzed particles per bin (red). Uncertainty analyses are given in the Supplement Sect. 7.

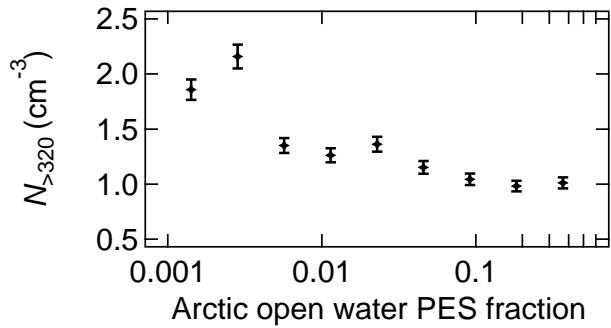


**Figure S23.** Aerosol composition measured by the ALABAMA as a function of FLEXPART derived air mass residence time over (a) vegetation fires, (b) populated areas, and (c) industrial areas during the southern air mass period. Figures represent each the cumulative particle fraction of nss-nitrate-containing (blue) and EC-containing (black) particles as well as the total number of analyzed particles per bin (red). Uncertainty analyses are given in the Supplement Sect. 7.

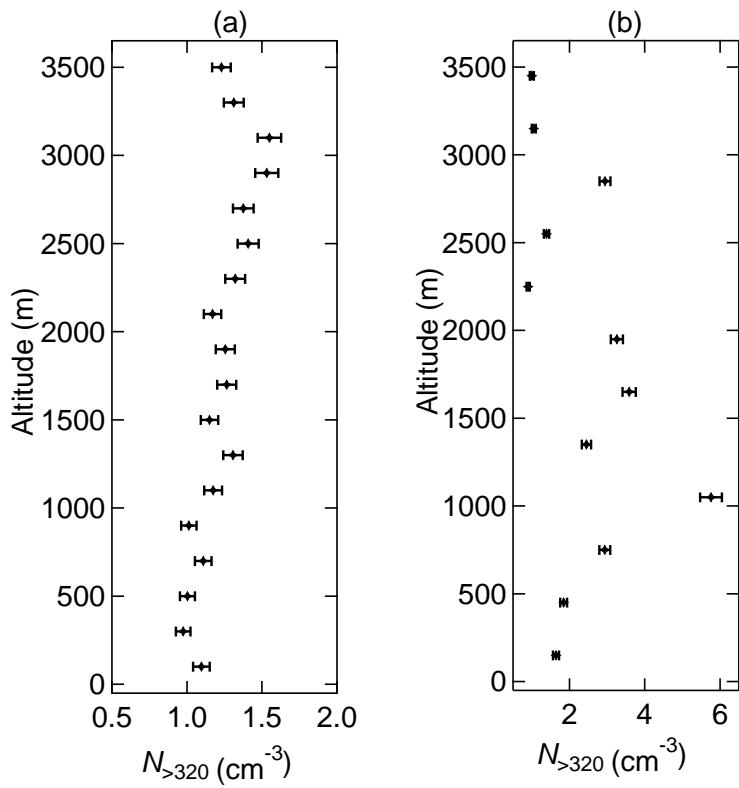


**Figure S24.** Aerosol composition measured by the ALABAMA as a function of FLEXPART derived air mass residence time over (a) populated areas and (b) industrial areas during the Arctic air mass period. Figures represent each the cumulative particle fraction of nss-nitrate-containing (blue) and EC-containing (black) particles as well as the total number of analyzed particles per bin (red). Uncertainty analyses are given in the Supplement Sect. 7.

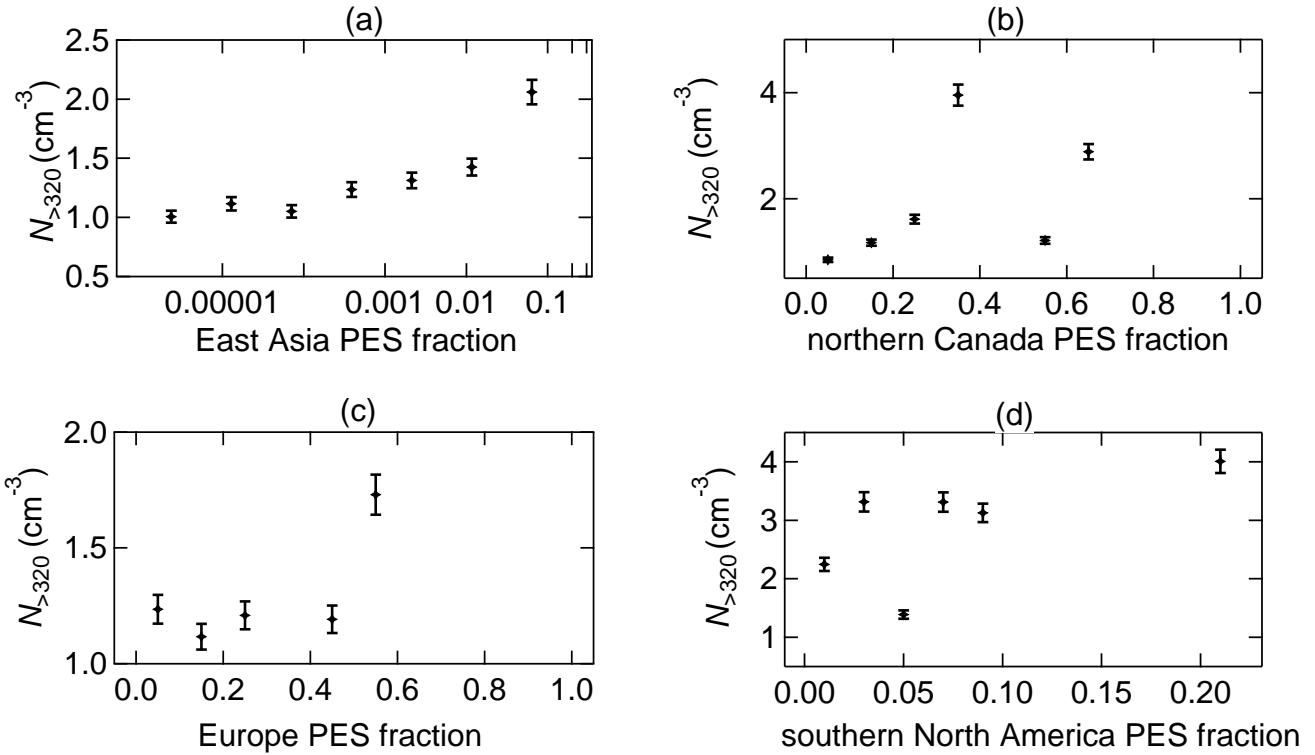
Figures S25-S32 provide the comparison between the number concentration of particles larger than 320 nm ( $N_{>320}$ ) measured by the UHSAS and various parameters.



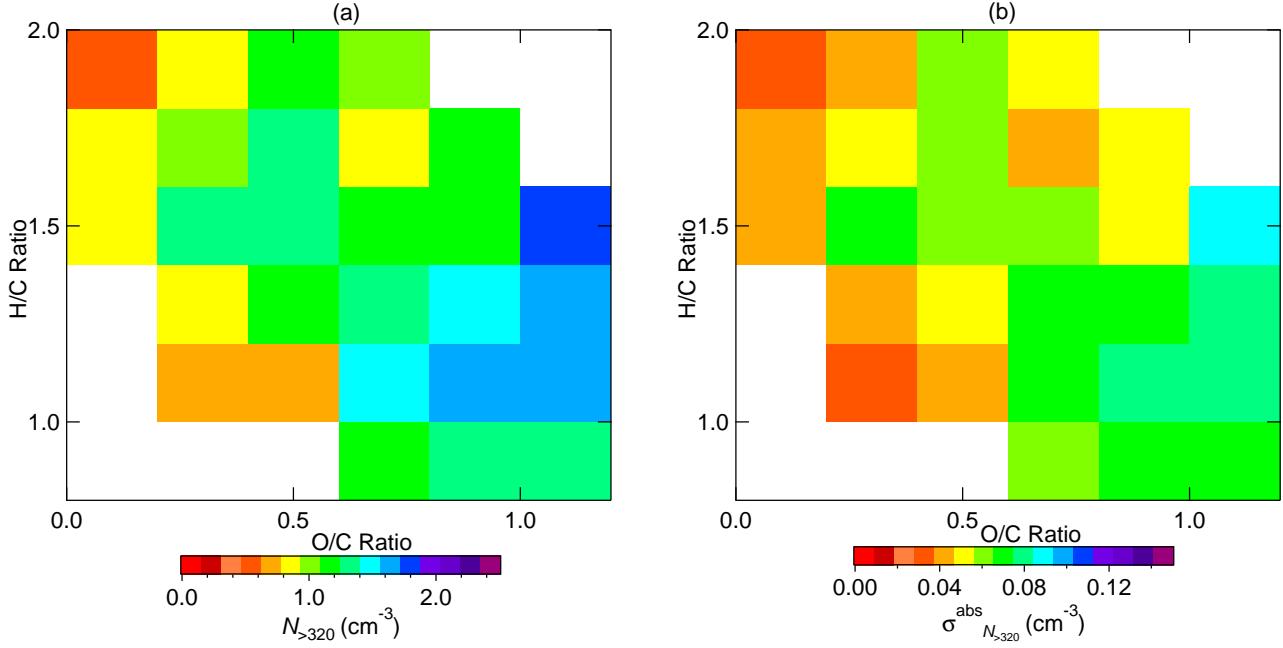
**Figure S25.** Averaged UHSAS number concentration in a size range between 320 and 870 nm ( $N_{>320}$ ) as a function of FLEXPART derived air mass residence time over Arctic open water areas during the Arctic air mass period. Uncertainty analysis is given in Sect. 7.



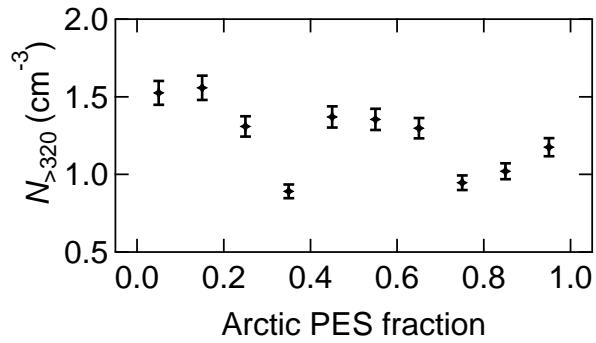
**Figure S26.** Vertically resolved averaged UHSAS number concentration in a size range between 320 and 870 nm ( $N_{>320}$ ) during the Arctic (a) and the southern (b) air mass periods. Uncertainty analysis is given in Sect. 7.



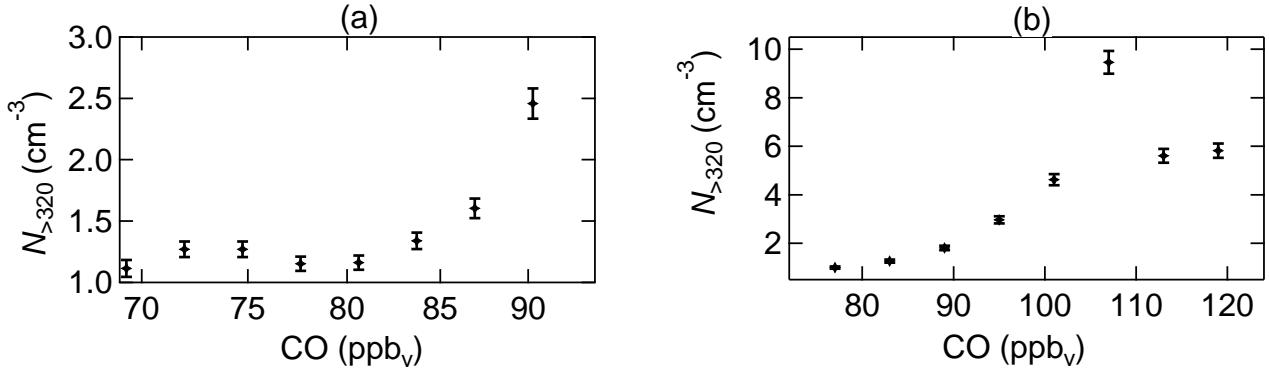
**Figure S27.** Averaged UHSAS number concentration in a size range between 320 and 870 nm ( $N_{>320}$ ) as a function of FLEXPART derived air mass residence time over the East Asia (a), northern Canada (b), Europe (c), and southern North America (d) during the Arctic (a, c) and the southern (b, d) air mass periods. Uncertainty analysis is given in Sect. 7.



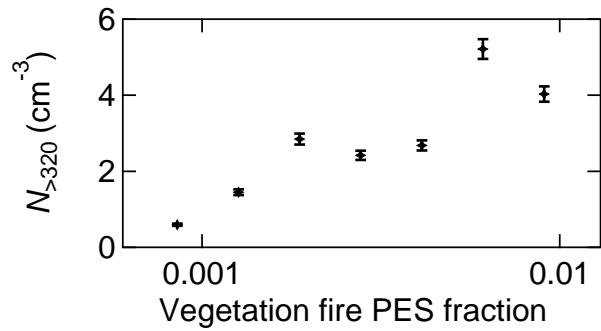
**Figure S28.** Comparison between the HR-ToF-AMS estimated oxygen-to-carbon (O/C) and hydrogen-to-carbon (H/C) ratios colored by (a) the averaged UHSAS number concentration in a size range between 320 and 870 nm ( $N_{>320}$ ) and (b) the absolute uncertainty of  $N_{>320}$  ( $\sigma_{N_{>320}}^{\text{abs}}$ ). Uncertainty analysis is given in Sect. 7.



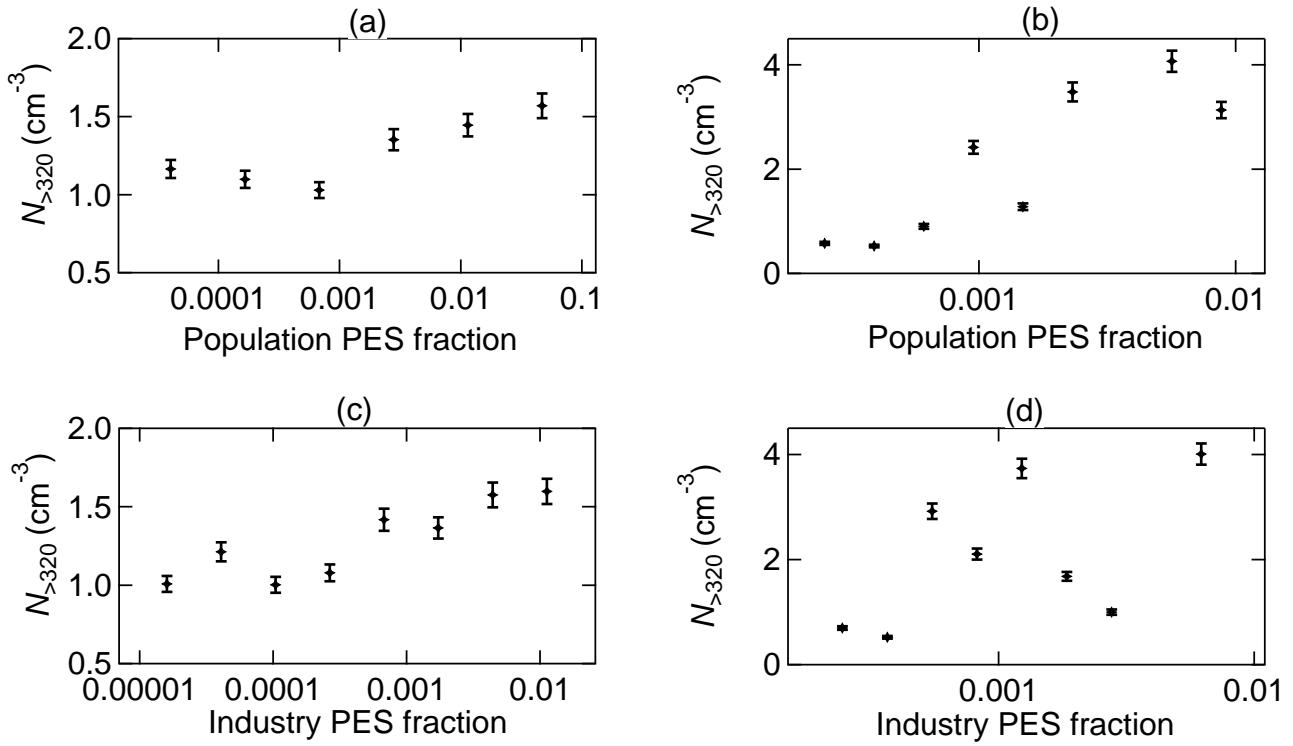
**Figure S29.** Averaged UHSAS number concentration in a size range between 320 and 870 nm ( $N_{>320}$ ) as a function of FLEXPART derived air mass residence time over the Arctic during the Arctic air mass period. Uncertainty analysis is given in Sect. 7.



**Figure S30.** Averaged UHSAS number concentration in a size range between 320 and 870 nm ( $N_{>320}$ ) as a function of CO mixing ratios during the Arctic (a) and the southern (b) air mass periods. Uncertainty analysis is given in Sect. 7.



**Figure S31.** Averaged UHSAS number concentration in a size range between 320 and 870 nm ( $N_{>320}$ ) as a function of FLEXPART derived air mass residence time over vegetation fires during the southern air mass period. Uncertainty analysis is given in Sect. 7.



**Figure S32.** Averaged UHSAS number concentration in a size range between 320 and 870 nm ( $N_{>320}$ ) as a function of FLEXPART derived air mass residence time over populated areas (a-b) and industrial areas (c-d) during the Arctic (a, c) and the southern (b, d) air mass periods. Uncertainty analysis is given in Sect. 7.

## 7 Uncertainty analyses

### 7.1 UHSAS number concentration ( $N_{>320}$ )

The absolute uncertainty of the averaged UHSAS number concentration for each bin ( $\sigma_{N_{>320}}^{\text{abs}}$ ) is calculated based on Poisson statistics and Gaussian propagation of uncertainties:

$$\sigma_{N_{>320}}^{\text{abs}} = N_{>320} \sqrt{\left(\sigma_{\text{UHSAS}}^{\text{rel}}\right)^2 + \left(\sigma_{\text{UHSAS}}^{\text{rel}_1}\right)^2}, \quad (3)$$

with the  $\sigma_{\text{UHSAS}}^{\text{rel}}$  that is defined as follows:

$$\sigma_{\text{UHSAS}}^{\text{rel}} = \frac{1}{\sqrt{C_{>320}}}, \quad (4)$$

with the counts of particles greater than 320 nm in diameter ( $C_{>320}$ ). The  $\sigma_{\text{UHSAS}}^{\text{rel}_1}$  is 5 % derived from the UHSAS collection efficiency of 95 % (see Table S3).

### 7.2 ALABAMA particle fraction ( $PF$ )

The absolute uncertainty of the ALABAMA particle fraction for each bin ( $\sigma_{PF}^{\text{abs}}$ ) is calculated using binomial statistics:

$$\sigma_{PF}^{\text{abs}} = \frac{\sqrt{N_{\text{hits}} \cdot PF \cdot (1 - PF)}}{N_{\text{hits}}}, \quad (5)$$

with the number of particles that are successfully ionized by the ablation laser and create a mass spectrum ( $N_{\text{hits}}$ ) and the  $PF$ .

### 7.3 ALABAMA scaled number concentration ( $PF \cdot N_{>320}$ )

The absolute uncertainty of the ALABAMA scaled number concentration for each bin ( $\sigma_{PF \cdot N_{>320}}^{\text{abs}}$ ) is calculated using Gaussian propagation of uncertainties:

$$\sigma_{PF \cdot N_{>320}}^{\text{abs}} = \sqrt{\left(PF \cdot \sigma_{N_{>320}}^{\text{abs}}\right)^2 + \left(N_{>320} \cdot \sigma_{PF}^{\text{abs}}\right)^2}. \quad (6)$$

### 7.4 Further information on complementary data

Table S3 presents further information on the instruments deployed in this study.

**Table S3.** Data and associated instrumentation used in this study along with additional information.

Instrument	Data used in this study [unit]	Particle size range <sup>a</sup>	Time resolution	T: total uncertainty	Ref.	Operator
UHSAS	$N_{>320}$ [ $\text{cm}^{-3}$ (STP) <sup>b</sup> ]	320 nm - 870 nm	1 s	CE: 95 % R: reproducibility A: accuracy	<sup>c</sup>	AWI
HR-ToF-AMS	ammonium, sulfate, nitrate, OM, MSA [ $\mu\text{g m}^{-3}$ (stdP) <sup>e</sup> ]	75 - 900 nm	20 - 30 s	DL: 0.004, 0.009, 0.008, 0.08, 0.005 $\mu\text{g m}^{-3}$ T: 33 - 37 %	<sup>d</sup>	UoT
SP2	rBC [ $\mu\text{g m}^{-3}$ (STP) <sup>b</sup> ]	85 - 700 nm	1 s	R: 10 %	<sup>f</sup>	AWI
ALABAMA	Na/Cl/Nitrate, EC, nss-nitrate, TMA, DCA [ $\text{cm}^{-3}$ (STP) <sup>h</sup> ]	320 - 870 nm	non-equidistant time steps max. 8 Hz	see Sect. S7	<sup>g</sup>	MPI-C
CO analyzer	CO [ppbv]		1 s	T: 4.7 ppbv	<sup>i</sup>	IPA

<sup>a</sup>The upper particle size cut-off is approximately 1  $\mu\text{m}$  owing to the particle transmission efficiency of the main aerosol inlet (Leaitch et al., 2016), although some instruments might have a higher upper cut-off.

<sup>b</sup>Data by the UHSAS and SP2 are presented at STP 273.15 K and 1013.25 hPa (Schulz et al., 2019)

<sup>c</sup>Cai et al. (2008); Schulz et al. (2019)

<sup>d</sup>DeCarlo et al. (2006); Bahreini et al. (2008); Willis et al. (2016, 2017)

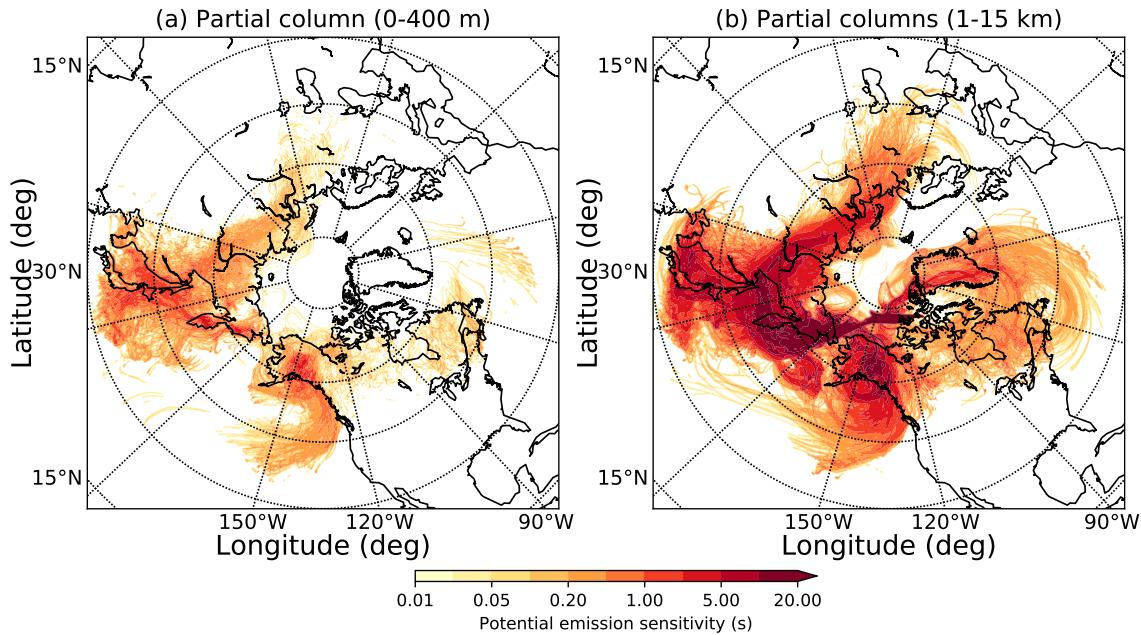
<sup>e</sup>Data by the HR-ToF-AMS are presented at standard pressure (stdP) 1013.25 hPa, since no temperature measurements inside the inlet line are available.

<sup>f</sup>Schwarz et al. (2006); Gao et al. (2007); Laborde et al. (2012); Dahlkötter et al. (2014); Schulz et al. (2019)

<sup>g</sup>Brands et al. (2011); Roth et al. (2016); Köhlher et al. (2017)

<sup>h</sup>Data by the ALABAMA are scaled to number concentrations using the UHSAS data that are presented at STP 273.15 K and 1013.25 hPa (Schulz et al., 2019)

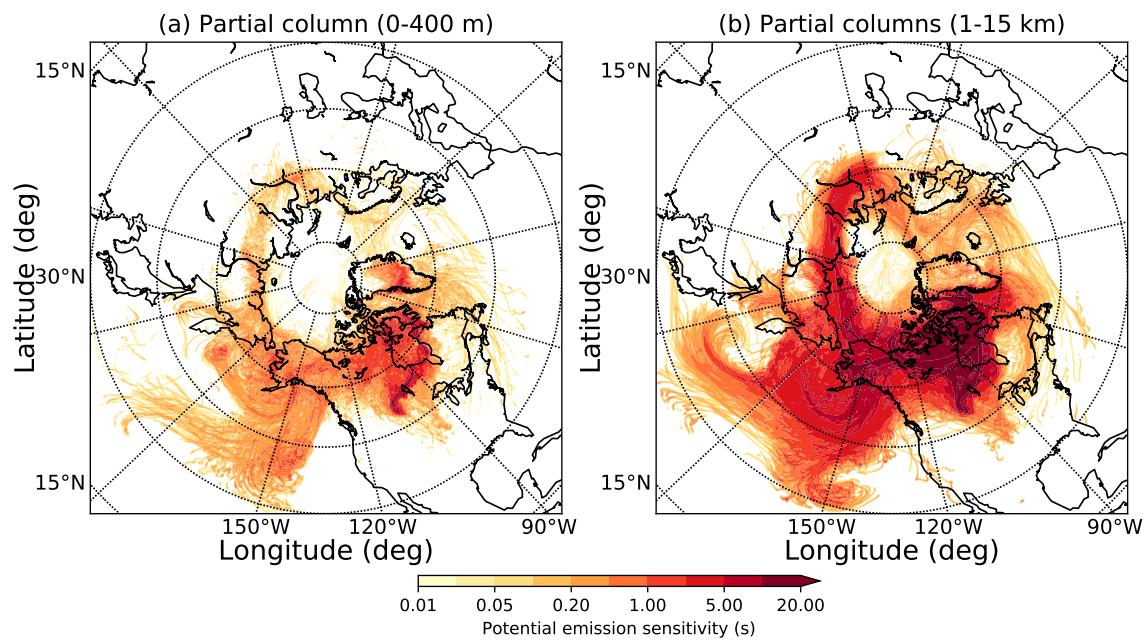
<sup>i</sup>Scharffe et al. (2012); Aero-Laser GmbH (2013); Bozem et al. (2019)



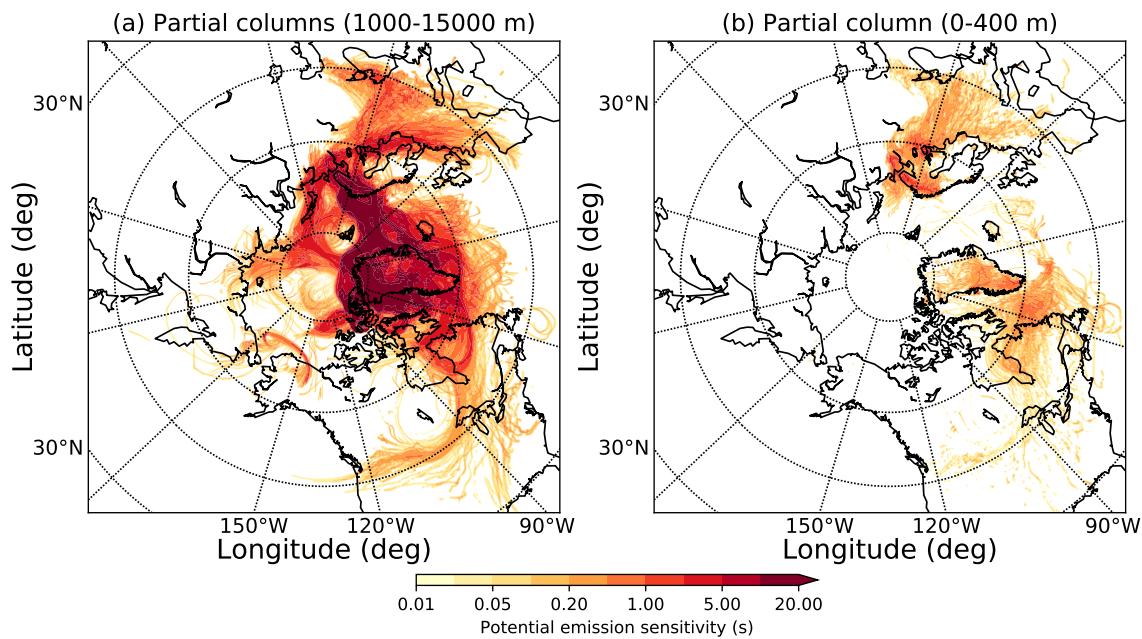
**Figure S33.** FLEXPART 15-days integrated (a) partial column 0 - 400 m and (b) partial column 1 - 15 km potential emission sensitivity (PES; for measurements on 5 July from 19:00 until 19:10 UTC) as an example of transport from sources in East Asia (particularly Japan) to the measurement region.

## 8 Examples of transport pathways

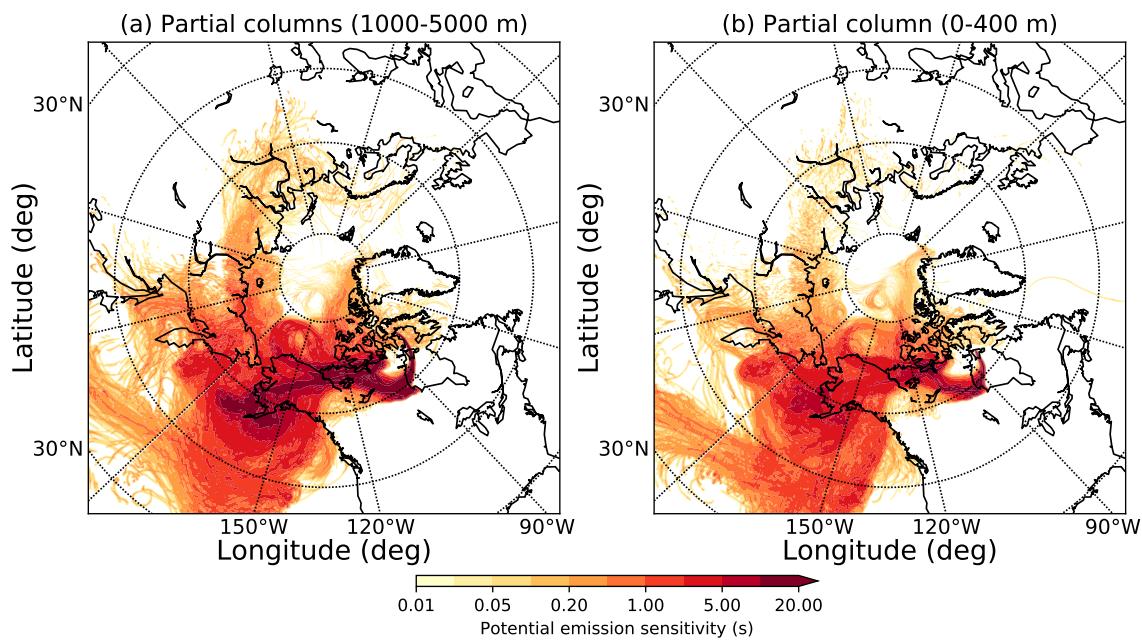
Figures S33-S36 provide examples of transport pathways from East Asia, southern North America, Europe, and northern Canada to Arctic regions, respectively.



**Figure S34.** FLEXPART 15-days integrated (a) partial column 0 - 400 m and (b) partial column 1 - 15 km potential emission sensitivity (PES; for measurements on 21 July from 18:30 until 18:40 UTC) as an example of transport from oil field emissions in Alberta (Canada) to the measurement region.



**Figure S35.** FLEXPART 15-days integrated (a) partial column 1 - 15 km and (b) partial column 0 - 400 m potential emission sensitivity (PES; for measurements on 12 July from 18:40 until 18:50 UTC) as an example of transport from European sources to the measurement region. Note that air mass transport into the Arctic took place mainly above 400 m.



**Figure S36.** FLEXPART 15-days integrated (a) partial column 1 - 5 km and (b) partial column 0 - 400 m potential emission sensitivity (PES; for measurements on 19 July from 15:30 until 15:40 UTC) as an example of transport of emissions from vegetation fires in northern Canada (compare to Fig. 4a) to the measurement region.

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