



1 **Observations and explicit modeling of isoprene chemical**
2 **processing in polluted air masses in rural areas of the Yangtze**
3 **River Delta region: radical cycling and formation of ozone and**
4 **formaldehyde**
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14

15 **Abstract**

16 Ozone pollution has become one of the most severe environmental problems in China in
17 recent years. Our online observations showed that high level of O₃ were observed in rural
18 areas of the Yangtze River Delta (YRD) region even there was no obvious ozone transport
19 from the urban regions. To better understand the formation mechanism of local O₃ pollution
20 and investigate the potential role of isoprene chemistry in the budgets of RO_x (OH+HO₂+RO₂)
21 radicals, synchronous observations of volatile organic compounds (VOCs), formaldehyde
22 (HCHO) and meteorological parameters were conducted at a rural site of the YRD region in
23 2018. Five episodes with elevated O₃ concentrations under stagnant meteorological
24 conditions were first identified; an observation-based model (OBM) with the Master
25 Chemical Mechanism was applied to analyze the photochemical processes in these high-O₃



26 episodes. High levels of O₃, nitrogen oxides (NO_x), and VOCs facilitated strong production
27 and recycling of RO_x radicals with the photolysis of oxygenated VOCs (OVOCs) being the
28 primary source. Our results suggest that local biogenic isoprene is important to local
29 photochemical processes. Removing isoprene could drastically slow down the efficiency of
30 RO_x recycling and reduce the concentrations of RO_x. The absence of isoprene chemistry
31 could further lead to decrease in the daily average concentration of O₃ and HCHO by 36%
32 and 15%, respectively. This study underlines that the isoprene chemistry in rural atmosphere
33 becomes important with the participation of anthropogenic NO_x and also provides insights
34 into the radical chemistry that essentially drives the formation of secondary pollutants (e.g.
35 O₃ and HCHO) in rural YRD region.

36 **Keywords:** Isoprene; Observation-based model (OBM); Radical; Ozone; Yangtze River
37 Delta

38 1. Introduction

39 The hydroxyl radical (OH), hydro peroxy radical (HO₂) and organic peroxy radical
40 (RO₂), collectively known as RO_x dominate the oxidative capacity of the atmosphere and
41 hence govern the removal of primary contaminants (e.g. volatile organic compounds (VOCs))
42 and the formation of secondary pollutants (e.g. ozone (O₃), secondary organic aerosols
43 (SOAs)) (Liu et al., 2012; Xue et al., 2016a). RO_x radicals can undergo efficient recycling (e.g.
44 OH→RO₂→RO→HO₂→OH) and produce O₃ and oxygenated VOCs (OVOCs) (Liu et al.,
45 2012; Tan et al., 2019; Xue et al., 2016b). In addition, the photolysis of OVOCs can in turn
46 produce primary RO₂ and HO₂ radicals, and further accelerate the recycling of RO_x (Liu et al.,
47 2012). The reaction rates of different VOCs with RO_x vary significantly (Atkinson and Arey,
48 2003; Atkinson et al., 2006). For instance, the reaction rate constants for OH with ethane and
49 ethene are 0.248×10^{-12} (cm molecule⁻¹ s⁻¹) and 8.52×10^{-12} (cm molecule⁻¹ s⁻¹), respectively.
50 Among the hundreds thousands of VOC species, isoprene (C₅H₈, 2-methyl-1,3-butadiene) is



51 one of the most active species, and also the most abundant biogenic VOCs (BVOCs) species
52 globally (Wennberg et al., 2018). Isoprene emissions from biogenic sources have been
53 extensively studied over past decades (Gong et al., 2018) and recent works have switched
54 from emissions to the degradation pathways and the impact of isoprene chemistry on regional
55 forest chemistry (Gong et al., 2018; Wolfe et al., 2016a). Previous studies showed that
56 isoprene could be quickly oxidized by atmospheric oxidants (e.g. OH, O₃ or NO₃) (Wolfe et
57 al., 2016a; Gong et al., 2018; Jenkin et al., 2015). Due to the rapid reaction between OH and
58 isoprene ($100 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ at 298 K), more than 90% of the total daytime
59 isoprene is removed via this reaction (Wennberg et al., 2018). The reaction between OH and
60 isoprene is initiated by the addition of OH and can generate isoprene hydroxyperoxy radicals
61 (ISOPO₂) (Wennberg et al., 2018; D'Ambro et al., 2017; Liu et al., 2013; Jenkin et al., 2015).
62 ISOPO₂ isomers could then interconvert rapidly due to reversible O₂ addition and are finally
63 removed via reactions with HO₂ or NO (Jenkin et al., 2015; Wolfe et al., 2016a). Hence, the
64 degradation process of isoprene is tightly associated with RO_x recycling. According to He et
65 al. (2019), isoprene chemistry could strongly influence the photochemical formation of O₃,
66 with a relative incremental reactivity (RIR) of ~0.06%/%. In addition to O₃, HCHO is formed
67 via several pathways during the depletion of isoprene (Jenkin et al., 2015; Wolfe et al., 2016a)
68 and HCHO formation is found to be highly sensitive to isoprene (Zeng et al., 2019).

69 The Yangtze River Delta (YRD) region is one of the most developed city-clusters in
70 eastern China and has been suffering from serious O₃ pollution (Zhang et al., 2019; Zhang et
71 al., 2020a; Chan et al., 2017). At the suburban area of YRD, high levels of O₃ were frequently
72 observed (Zhang et al., 2019; Zhang et al., 2020a). A number of studies have been conducted
73 to investigate the relationships between O₃ precursors and O₃ (Chan et al., 2017; Lin et al.,
74 2020; Zhang et al., 2020a; Zhang et al., 2020b), but few have attempted to address the
75 atmospheric oxidizing capacity and radical chemistry involved in these complicated



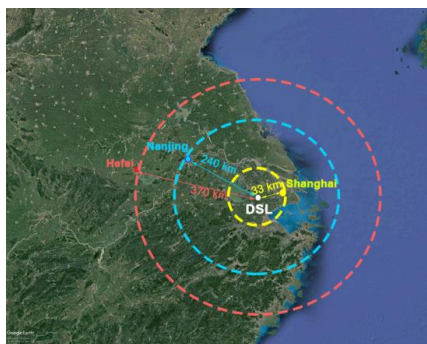
76 photochemical processes (Tan et al., 2019;Zhu et al., 2020b). Previous studies have pointed
77 out that high level of O₃ at suburban areas of Shanghai could be attributed to the transport of
78 O₃ or O₃ precursors from urban areas (Lin et al., 2020; Zhang et al., 2019). However, high O₃
79 concentration was frequently observed in suburban areas under stable meteorological
80 conditions. Given the high vegetation coverage in rural YRD and weak transport of air
81 masses, the importance of local isoprene chemistry to ozone formation remains unclear.

82 In this study, we conducted a comprehensive set of in-situ observations of isoprene,
83 meteorological parameters, and trace gases to understand the importance of isoprene
84 chemistry to atmospheric photochemical processes in rural YRD region. An observation-
85 based model (OBM) was used to explore the role of local isoprene chemistry in radical
86 budgets and the formation of O₃ and HCHO. Results from this study can provide insights of
87 the isoprene chemistry in the rural region of a fast-developing city-cluster.

88 **2. Methodology**

89 **2.1 Measurement site and techniques**

90 The observations were conducted at a supersite (120.98°E, 31.09°N) in the rural areas of
91 the YRD region (Figure 1). It is located to the west of Shanghai and is close to the Dianshan
92 Lake Scenic area, which has high vegetation coverage. To investigate the local isoprene
93 chemistry and its influence on O₃ and HCHO formation, continuous measurements were
94 conducted from Apr. 7 to Sep. 25, 2018, when the photochemical reactions are active and
95 ozone formation is significant.



96

97 **Figure 1. Location of the Dianshan lake supersite (white dot) in the rural areas of YRD region. This**
98 **picture was created with © Google Earth on 23 July 2020.**

99 **Table 1. Measurements performed during the ozone season**

Species/Parameter	Experimental Technique	Time resolution
O ₃	Model 49i, Thermo Fischer Scientific, USA	10 s
NO and NO ₂	Model 43i, Thermo Fischer Scientific, USA	60 s
SO ₂	Model 42i, Thermo Fischer Scientific, USA	80 s
HCHO	AL4021, Aero-Laser, GER	90 s
VOCs species	GC866, Agilent., USA	1 hour
Temperature, relative humidity, wind speed and wind direction	Meteorological station, Vaisala, NLD	60 s

100

101 The measurements performed are shown in Table 1. Wind speed (WS), wind direction
102 (WD), temperature (T), and relative humidity (RH) were simultaneously observed by a
103 meteorological station (Vaisala, FIN). According to China's air quality standard, several
104 criteria air pollutants were measured during this experiment. O₃ was measured by an
105 ultraviolet photometric analyzer (Model 49i, Thermo Fischer Scientific., USA), which has a
106 detection limit of 10 ppbv at 10 second resolution. 1 min resolution of nitrogen oxides (NO
107 and NO₂) data were simultaneously observed by a chemiluminescence instrument (Model 43i,
108 Thermo Fischer Scientific., USA), which has a detection limit of 0.40 ppbv. Sulfur dioxide
109 was monitored by a pulse fluorescence analyzer (Model 42i, Thermo Fischer Scientific., USA)
110 with a detection limit of 0.50 ppbv at 300 second resolution. Carbon monoxide was
111 monitored by a gas filter correlation infrared absorption analyzer (Model 48i, Thermo Fischer



112 Scientific., USA), which has a detection limit of 0.04 ppm. All the online instruments used
113 for gas analyzer were auto-zero every day, and were multi-point calibrated every month. All
114 the instruments used for the online observation were housed on top of a 5-floor-high building,
115 about 15 m above the ground level.

116 A total of 55 VOC species, including 28 alkanes, 10 alkenes, 16 aromatics and acetylene
117 were continuously analyzed at our sampling site by two online gas chromatograph with flame
118 ionization detector (GC-FID) systems (GC-866 airmoVOC C₂-C₆ #58850712 and airmoVOC
119 C₆-C₁₂ #283607112, Agilent., USA) with a time resolution of 1 hour during our experiment.
120 Ambient samples are directly inhaled into this system by a pump. Low carbon VOCs (C₂-C₆)
121 are captured by a low temperature (-10 °C) preconcentration system, while high carbon
122 VOCs are concentrated by a built-in room temperature preconcentration system. Then the
123 preconcentration systems are heated and desorb VOCs, which are then carried into
124 chromatographic columns by helium. Individual VOCs separated in the columns are
125 eventually detected by FID systems. Formaldehyde (HCHO) was continuously measured by a
126 Hantzsch fluorescence technique (AL4201, Aerolaser GmbH., GER), which is based on
127 fluorometric Hantzsch reaction in the liquid phase requires the quantitative transfer of HCHO
128 from gas phase to liquid phase. A Hantzsch reagent (acetylacetone) was used in this
129 instrument.

130 **2.2 Observation-based model**

131 A user-friendly zero-dimensional (0-D) box model (F0AM) was used to simulate the
132 chemical processes in the atmosphere in this study. This model was developed by Wolfe et al.
133 (2016b) based on University of Washington Chemical Model (UWCM). Dry deposition, aloft
134 exchange and atmospheric dilution were considered in this model. We chose the Master
135 Chemical Mechanism (MCM) v3.3.1 as the chemical mechanism with more than 5900
136 chemical species and 16500 reactions, which enables a detailed description of the complex



137 reactions. Photolysis frequencies (J values) were calculated by a trigonometric
138 parameterization based on solar zenith angle (SZA):

$$J = I \cos(SZA)^m \exp(-n \sec(SZA)) \quad (1)$$

139 where I , m and n are constants unique to each photolysis reaction, derived from least-squares
140 fits to J values computed with fixed solar spectra and literature cross-section and quantum
141 yields (Wolfe et al., 2016b). Hourly averaged concentrations of speciated VOCs, O₃, NO,
142 NO₂, SO₂, and meteorological parameters (such as T, TH, P) were used to constrain the
143 F0AM model. Nitrous acid (HONO) was not measured during our observation. Therefore, it
144 was fixed as 2% of the observed NO₂ concentration. This constant ratio is well observed in
145 different field studies and performed well in previous box model studies (Tan et al., 2019).
146 Before each simulation, the model will run 3 days as spin up to reach a steady state for
147 unmeasured species (e.g., OH and NO₃ radicals). The comparison of simulated observed O₃
148 concentrations can be shown in Figure S1 (Supplemental Information). The index of
149 agreement (IOA), mean bias (MB) and normalized mean bias (NMB) of O₃ in this simulation
150 were 0.90, 0.76 and 0.10, respectively. This result suggests that the model can reasonably
151 reproduce the variations of O₃ and could be used for further analysis. To quantify the
152 atmospheric oxidative capacity (AOC) changes in response to isoprene chemistry, two
153 parallel scenarios (S0 and S1) were conducted with isoprene chemistry disabled in S1. In
154 both cases, identical chemical mechanism and meteorological conditions were used to drive
155 model simulations. Through a comparative analysis of the scenarios, the impact of isoprene
156 chemistry on AOC can be obtained.

157 **3. Results and discussions**

158 **3.1 Overview of the observations**

159 To investigate the impact of local chemistry on ozone formation, five days with low
160 daily average wind speed (<2m/s) and high maximum daily 8-h average (MDA8) O₃



161 concentration (>74.7 ppb) were identified as typical local chemistry cases. Figure 2 shows the
162 time series of observed meteorological parameters (P, T, and RH), trace gases (NO, NO₂ and
163 O₃), isoprene and HCHO on selected days. On those selected days, the air masses reaching
164 the site were mainly from southeast and southwest (Figure 2). The weak wind was not
165 conducive to the regional transportation of air pollutants. The observed O₃, NO₂, NO and CO
166 ranged from 1.40 to 155.40 ppbv (52.72 ± 44.43 ppbv, average value), 5.36 to 57.95 ppbv
167 (21.58 ± 12.88 ppbv), 0.75 to 54.51 ppbv (5.40 ± 8.13 ppbv), and 400 to 960 ppbv (597 ± 153
168 ppbv), respectively. The conditional probability function (CPF) is applied to exhibit the
169 relationship between high O₃ concentrations and wind (Figure 3). The detailed description of
170 CPF can be found in supplemental information. The result suggests that pretty high O₃
171 concentrations (>131 ppb) was usually observed when the site was influenced by weak south
172 wind. This implies that high O₃ was most likely produced locally. Although this site is far
173 away from urban areas, high levels of NO were found during early morning, which is likely
174 caused by nearby fresh emissions. As for NO₂, only one peak was found at dusk. This was
175 different from the results in urban areas (Zhang et al., 2019). It is worth noting that NO₂ and
176 O₃ concentrations were high even during nighttime, suggesting that the AOC remained high
177 at nighttime. The daily average isoprene concentrations were 0.37 ± 0.36 ppbv, which is
178 comparable to that observed by Gong et al. (2018) at a forested mountaintop site ($0.287 \pm$
179 0.032 ppbv). The average HCHO was 5.01 ± 3.80 ppbv, which was ~ 2 times of that observed
180 at a rural site of Hong Kong (Yang et al., 2020). It is worth noting that HCHO could reach
181 18.69 ppbv at midday.
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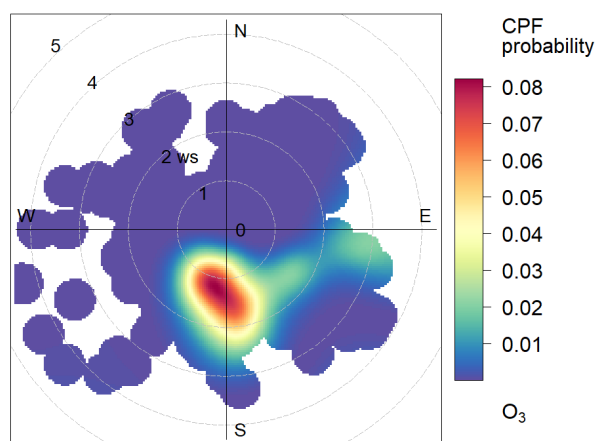


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Figure 2. Time series of hourly averages for trace gases, isoprene, HCHO, and meteorological parameters.

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CPF at the 95th percentile (=131)

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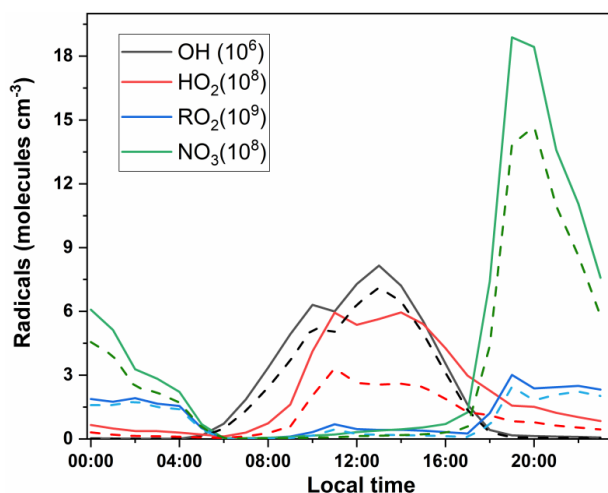
Figure 3. CPF polar plot of O₃ at DSL to

188 3.2 Simulated concentrations of radicals

189 Figure 4 shows the simulated average diurnal variation of major radicals in the base
190 scenario (S0). The average concentrations of OH, HO₂, RO₂, and NO₃ were estimated at 2.39



191 $\times 10^6$, 2.21×10^8 , 1.11×10^9 , and 4.24×10^8 molecules cm^{-3} , respectively. The simulated daily
192 average OH concentration lies between the simulated values during the summer in Beijing (9
193 $\times 10^6$ molecules cm^{-3}) and the simulated value at a suburban site in Hong Kong in 2013 ($1.5 \pm$
194 0.2×10^6 molecules cm^{-3}) (Liu et al., 2019; Xue et al., 2016a). In addition, the average
195 daytime OH concentration was $\sim 50\%$ lower than that simulated at a forested mountaintop site
196 in southern China (Gong et al., 2018). The maximum HO_2 concentration simulated for DSL
197 site (5.95×10^8 molecules cm^{-3}) was close to that reported in Wuhan (4.7×10^8 molecules cm^{-3})
198 (Zhu et al., 2020a), but was $\sim 13\%$ lower than that in Beijing (6.8×10^8 molecules cm^{-3}) (Liu et
199 al., 2012). Pretty high level of NO_3 (as high as $\sim 19 \times 10^8$ molecules cm^{-3}) was found during
200 nighttime. The average nocturnal NO_3 concentration was 8.15×10^8 molecule cm^{-3} , which
201 was $\sim 36\%$ higher than that simulated in the study of Gong et al. (2018). As aforementioned,
202 during nighttime, pretty high level of NO_2 (27.71 ppbv) and O_3 (30.05 ppbv) was observed,
203 which favored the formation of NO_3 . Interestingly, a high level of RO_2 was also found during
204 nighttime. This result is different from the study of Liu et al. (2012), which found the
205 maximum value of RO_2 during daytime. By separate the formation of RO_2 , we found that
206 during nighttime, over 70% RO_2 was formatted via the oxidation of RO by NO_3 radical,
207 suggesting that the nighttime chemistry in the suburban site was also very important.



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209 **Figure 4.** Average diurnal variation of OH, HO₂, RO₂ and NO₃ in S0 (solid lines) and S1 (dash lines).

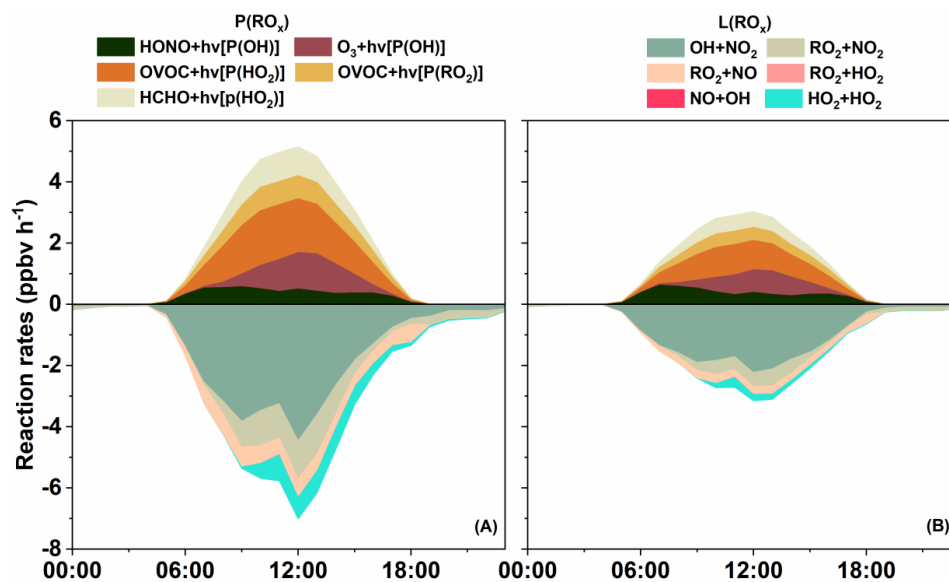
210 3.3 Recycling of RO_x radicals

211 Figure 5(A) shows the primary sources of RO_x in S0 and the detailed daytime budget of
212 RO_x. Minor RO_x sources, e.g. ozonolysis of alkenes, are not shown. Heterogeneous reactions
213 were not considered in this model. The photolysis of O₃ becomes the predominant primary
214 source of OH, with a daytime mean production rate of 0.52 ppbv h⁻¹, which was comparable
215 to that found by Liu et al. (2012) in Beijing, but was 0.38 ppbv h⁻¹ lower than the result in the
216 study of Xue et al. (2016). The photolysis of HONO was the second largest primary OH
217 source at the DSL site, contributing 0.41 ppbv h⁻¹ of daytime OH production in our
218 simulation. This value is much lower than the results of Liu et al. (2012 and 2019) and Xue et
219 al. (2016b). Such low value was most likely caused by the excessive constrain on HONO
220 since HONO was not directly monitored during our experiment. As mentioned in the study of
221 Liu et al. (2012), the photolysis processes of excess HONO from heterogeneous source can
222 be the predominant primary OH source. Unfortunately, we were not able to quantitatively
223 consider this important mechanism in this study due to the lack of relevant observation, and
224 further studies are needed to make up this limitation. As for HO₂, the photolysis of OVOC is
225 the predominant source with a daytime mean production rate of 1.09 ppbv h⁻¹ and maximum



226 reaching 1.79 ppbv h^{-1} , which is comparable to Xue et al. (2016). The photolysis of HCHO
227 can also contribute 0.54 ppbv h^{-1} to the daytime production of HO_2 , which is close to the
228 results of Xue et al. (2016). As for RO_2 , the photolysis of OVOC was the largest source (0.49
229 ppbv h^{-1}), which was relatively lower than the results found at urban site (Liu et al., 2012).
230 From the RO_x perspective, the daytime primary radical production in DSL site was
231 dominated by the photolysis of OVOCs (except for HCHO), followed by the photolysis of
232 HCHO and O_3 . Summing up all the sources of RO_x gives a total primary daytime RO_x
233 production rate of 3.06 ppbv h^{-1} (0.94 ppbv h^{-1} for OH, 1.63 ppbv h^{-1} for HO_2 , and 0.49 ppbv
234 h^{-1} for RO_2), which was 54~62% lower than those in Beijing (6.6 ppbv h^{-1}) (Liu et al., 2012)
235 and Hong Kong (8.11 ppbv h^{-1}) (Xue et al., 2016b).

236 RO_x radicals are ultimately removed from the atmosphere via deposition of radical
237 reservoir species, e.g. H_2O_2 , HNO_3 , and ROOH (Liu et al., 2012). The terminate processes of
238 RO_x was dominated by their reactions with NO_x . Specifically, the reaction of $\text{OH}+\text{NO}_2$,
239 RO_2+NO_2 , RO_2+NO , forming HNO_3 , RO_2NO_2 , and RONO_2 , accounting for 2.46, 0.63, and
240 0.53 ppbv h^{-1} of the RO_x radical loss during daytime. This is consistent with the
241 understanding that reactions with NO_x usually dominate the radical sink in high NO_x
242 environments (Xue et al., 2016b; Liu et al., 2012). In addition, RONO_2 and RO_2NO_2 could in
243 turn react with OH, leading to 0.45 ppbv h^{-1} of daytime OH sinks (Figure 6). Summing up the
244 primary sources and sinks gives a negative value of net RO_x production, suggesting that the
245 RO_x was in a stage of gradual depletion.



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Figure 5. Primary daytime sources of RO_x in S0 (A) and S1 (B).

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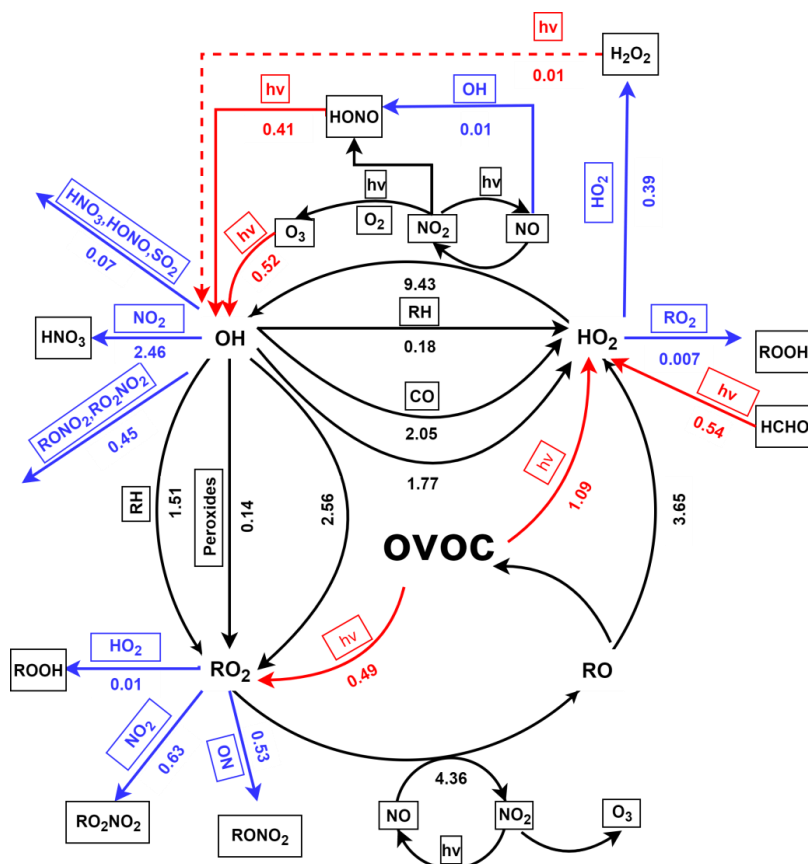
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The daytime (6:00-18:00) recycling of RO_x is shown in Figure 6, with primary sources (in red) and sinks (in blue) of RO_x. In the recycling of RO_x, the production of OH was dominated by the reaction of HO₂+NO (9.43 ppbv h⁻¹). As for RO₂, it was produced by the reaction of OH with OVOC (2.56 ppbv h⁻¹), alkyl (RH) (1.51 ppbv h⁻¹), and peroxides (0.14 ppbv h⁻¹). The reaction of RO₂+NO can result in strong production of RO (4.36 ppbv h⁻¹). The reaction of RO and O₂ was the major contributor to HO₂ production, followed by the reaction of OH with CO (2.05 ppbv h⁻¹), OVOC (1.77 ppbv h⁻¹), and RH (0.18 ppbv h⁻¹). It should be noted that the top two fast reactions within the recycling of RO_x (HO₂+NO and RO₂+NO) were related to NO_x. As mentioned in the study of Liu et al. (2012), this result could be mainly due to the abundance of NO (e.g. ~50 ppbv in the morning). Obviously, these recycling processes dominate the total production of OH, HO₂ and RO₂ radicals. As suggested in the study of Xue et al. (2016) and Liu et al. (2012), the radical propagation is efficient and enhances the effect of the newly produced radicals in the polluted atmospheres with the co-existence of abundant NO_x and VOCs.



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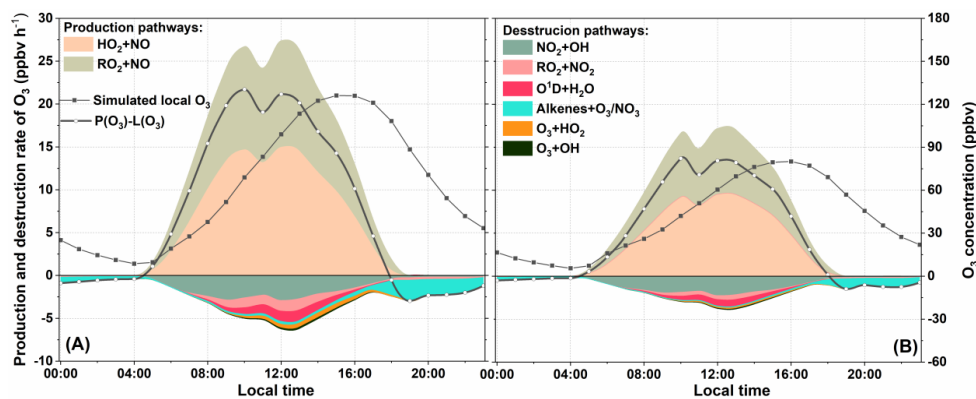
263 **Figure 6. Summary of daytime (06:00-18:00) average budgets of RO_x radicals (in ppbv h⁻¹). Primary RO_x**
 264 **sources and sinks are in red and blue.**

265 3.4 Production and destruction of O₃

266 Figure 7 illustrates the diurnal variation of simulated O₃ production and destruction
 267 pathways in S0. Also shown is the simulated average diurnal pattern of O₃ concentration and
 268 the net O₃ production rate. In the troposphere, the formation of O₃ is via the reactions of NO
 269 with peroxy radicals (e.g. HO₂ and RO₂) (Liu et al., 2012;Xue et al., 2016b;Zhu et al., 2020a).
 270 On average, the reaction of HO₂+NO and RO₂+NO attributed 5.10 and 4.37 ppbv h⁻¹ of the
 271 production of O₃. The maximum rate of HO₂+NO (14.96 ppbv h⁻¹) and RO₂+NO (12.42 ppbv
 272 h⁻¹) both occurred at 12:00. The total daytime production rate of O₃ (P(O₃)) is the sum of
 273 HO₂+NO and RO₂+NO at 17.48 ppbv h⁻¹, which lies between that simulated in Beijing (32



274 ppbv h⁻¹) (Liu et al. 2012) and Hong Kong (6.7 ppbv h⁻¹) (Liu et al., 2019). Due to the fast
275 cycling of both O₃ and NO₂, O₃ loss was due to several reactions leading to the destruction of
276 O₃ and NO₂. In our cases, the reaction of NO₂+OH becomes the predominant scavenging
277 pathways of O₃, with an average daytime reaction rate of 1.97 ppbv h⁻¹ (49.8%, percentage of
278 the total O₃ destruction rate, same below). This is comparable to the study of Liu et al. (2012
279 and 2019). The reaction of RO₂ and NO₂ was the second contributor to O₃ loss, with a mean
280 contribution of 0.63 ppbv h⁻¹ (16%). Other pathways, e.g. photolysis of O₃, ozonolysis of
281 alkenes, and O₃+HO₂, contributed 1.31 ppbv h⁻¹ of the total destruction rate of O₃ during
282 daytime. The daytime mean L(O₃) rate was 3.85 ppbv h⁻¹, which was ~22% of P(O₃),
283 suggesting that O₃ could efficiently accumulate during daytime. The net production of O₃
284 (P(O₃)-L(O₃)) is also shown in Figure 7. The maximum O₃ concentration was found at
285 around 16:00, which was also observed in other rural sites (Zong et al., 2018;Zhang et al.,
286 2019). It is worth noting that, the reaction of alkenes+O₃/NO₃ serves as an important pathway
287 of O₃ loss during nighttime (as high as 2.41 ppbv h⁻¹).



288
289 **Figure 7.** Average diurnal profiles of O₃ production and destruction rates (ppbv h⁻¹) in S0 (A) and S1 (B).

290 3.5 Production and destruction of HCHO

291 As aforementioned, high levels of HCHO was observed at DSL, a rural site over the
292 Yangtze River Delta region. Figure 8 (A) shows the production and loss pathways of HCHO



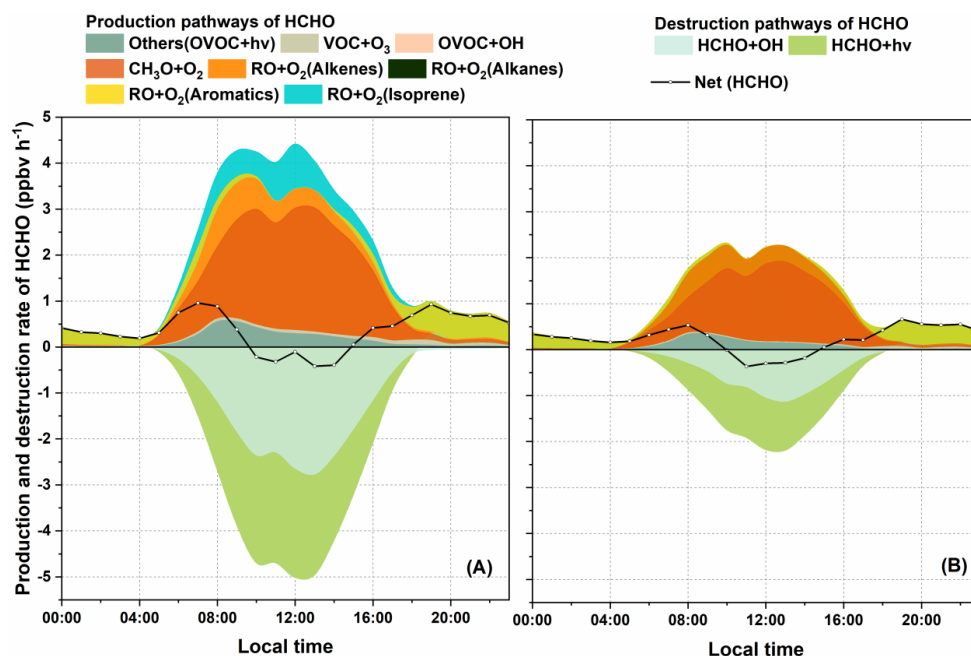
293 in S0. On average, HCHO production was dominated by the reaction of RO+O₂, accounting
294 for ~90% of the total production rate. Further classification of RO+O₂ pathway suggested that
295 the oxidation of CH₃O made a significant contribution of ~47%, followed by RO (from
296 isoprene) + O₂ reaction (12%) and RO (from aromatics) + O₂ reaction (~11%). This result is
297 comparable to the study of Yang et al. (2020 and 2018). Inspection of the model result shows
298 that CH₃O was mainly produced during the degradation of small molecular weight alkanes,
299 especially methane. It is notable that the reaction of RO (from aromatics) + O₂ could become
300 the predominant pathway of HCHO production during nighttime. This could be attributed to
301 the high level of NO₃ during nighttime, by which styrene could be quickly oxidized and
302 generate N-containing RO radicals, and furtherly generate HCHO. During daytime, isoprene
303 became the most important VOC specie of HCHO production, with a mean rate of 0.38 ppbv
304 h⁻¹. As mentioned, this site is surrounded by highly vegetated areas, which can provide
305 abundant biogenic isoprene. During daytime, over 90% of isoprene was oxidized by OH
306 radicals (Figure S2). According to MCMv3.3.1, several RO₂ species (e.g. ISOP34O₂,
307 ISOPDO₂, ISOPCO₂, CISOPAO₂, ISOPAO₂) can be generated during the OH-initiated
308 degradation process of isoprene (Jenkin et al., 2015). With the present of NO, isoprene-
309 originated RO₂ can transfer into RO (e.g. ISOPDO, ISOP34O, ISOPAO). The subsequent
310 degradation processes of isoprene-related RO, especially ISOP34O, ISOPDO, ISOPAO and
311 ISOPBO, are tightly related to the formation of HCHO (Jenkin et al., 2015). Other sources of
312 HCHO, such as the reaction between VOC and O₃, photolysis of OVOC and the reaction of
313 OVOC+OH only contributed minor amount of the total production rate during whole day.

314 As for HCHO depletion, the photolysis of HCHO and the reaction of HCHO+OH was
315 the two dominate pathway, accounting for ~50% and ~50% of the total depletion rate,
316 respectively. The net HCHO (equals to P(HCHO) + L(HCHO)) production rate was also
317 shown in Figure 8. After sunrise, the net production rate of HCHO raised gradually until 8:00,



318 when it reached the maximum rate (1.10 ppbv h^{-1}). This result is comparable to the study of
319 Yang et al. (2018). At around 10:00, the net(HCHO) dropped to $\sim 0 \text{ ppbv h}^{-1}$, that was roughly
320 consistent with our observation, which shows that the HCHO peak occurs at around 11:00.
321 After 10:00, although the reaction of $\text{RO}+\text{O}_2$ quickly produced HCHO, the depletion
322 pathways, especially the photolysis of HCHO, became more competitive, leading to the net
323 reduction of HCHO. After 12:00, the photolysis of HCHO dropped rapidly and the net
324 depletion of HCHO decreased to $\sim 0 \text{ ppbv h}^{-1}$ at around 15:00. The daytime net HCHO
325 production rate was 0.89 ppbv h^{-1} , which was comparable to result of Yang et al. (2018).

326 The above analysis indicates that the photolysis of OVOC, HCHO, O_3 and HONO was
327 the primary source of RO_x , which offers high oxidizing environment for the degradation of
328 VOCs. As a typical by-product in the degradation of several VOCs, HCHO can be quickly
329 formatted during daytime. The insight into detailed photochemical processes shows the
330 important role of isoprene in the formation of HCHO.





332 **Figure 8. Average diurnal profiles of net rate (net (HCHO)), breakdown HCHO production rate and**
333 **destruction rate (ppbv h⁻¹) in S0 (A) and S1 (B).**

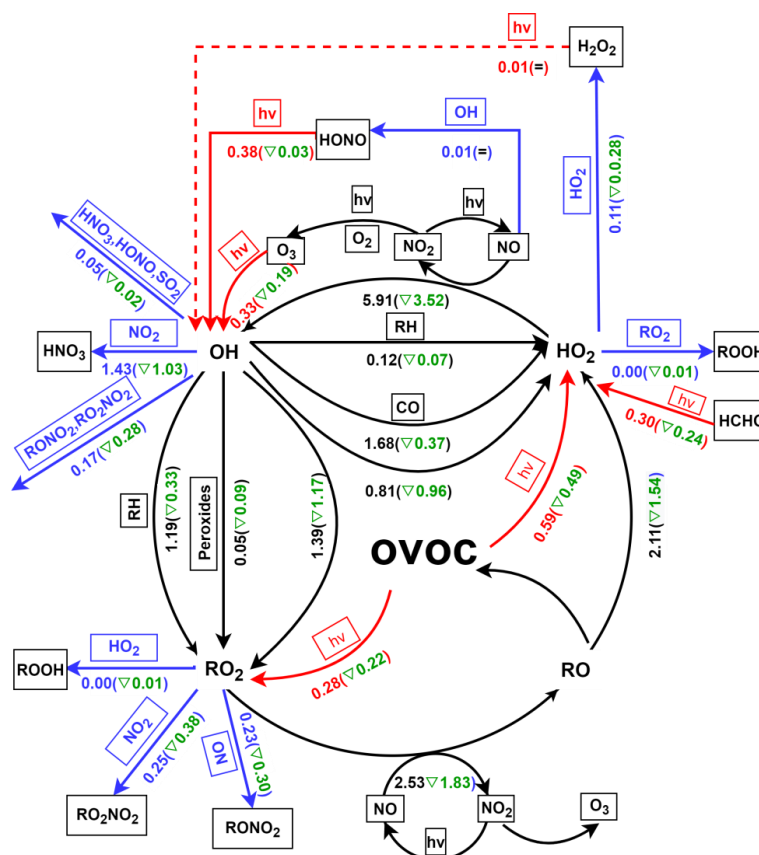
334 **3.6 Impacts of isoprene chemistry on photochemistry**

335 3.6.1 Impact on RO_x budget

336 As aforementioned, the degradation of isoprene is tightly related to the cycling of RO_x.
337 To roughly explain the impact of isoprene chemistry on RO_x budget, we carried out a parallel
338 simulation (S1) where isoprene chemistry is disabled (see in Figure 9). The diurnal variation
339 of OH, HO₂, RO₂ and NO₃ in S1 is also shown in Figure 4 (B). Most of the reaction rates
340 show a decrease trend in S1, suggesting that the absence of isoprene slows down the RO_x
341 recycling. The photolysis of OVOC (0.87 ppbv h⁻¹) is still the predominant primary source of
342 RO_x. However, without isoprene, the photolysis rate of OVOC decreased by 0.71 ppbv h⁻¹.
343 The total production and depletion rate of OH dropped to 6.22 and 5.53 ppbv h⁻¹, respectively.
344 Although the absence of isoprene could reduce the consumption of OH, the OH concentration
345 would be reduced by ~16% compared to S0, suggesting that the amount of OH produced via
346 isoprene chemistry is large enough to compensate for the shift from OH to peroxy radicals in
347 the RO_x family. As for RO₂, the daytime production and destruction rate falls to 1.87 and
348 3.01 ppbv h⁻¹, respectively. This means the concentration of RO₂ would be in a stage of
349 gradual decline. In addition, the absence of isoprene could also reduce RO₂ concentration by
350 ~20%, suggesting that isoprene was an important source of RO₂ at DSL site. As for HO₂,
351 drastic decrease of ~53% was found in S1. The above-mentioned decrease in RO_x obviously
352 could not be explained solely by the addition of isoprene-related radicals. Inspection of the
353 model results shows that OVOCs concentrations decreased drastically (~40%) after cutting
354 isoprene (e.g. ~44% decrease in formaldehyde, ~60% decrease in methylglyoxal, ~46%
355 decrease in glyoxal, ~100% decrease in methacrolein (MACR), and ~100% decrease in
356 methyl vinyl ketone (MVK)). The decrease in OVOC can further pull down substantial
357 amount of primary RO₂ and HO₂ (Figure 6 and Figure 9). It is interesting to note that,



358 subtracting isoprene also cause drop of NO_3 (~11%). This result can be contributed to the
 359 decrease of secondary production of O_3 (~36%), which can further reduce the formation of
 360 NO_3 , especially during nighttime.



361
 362 **Figure 9. Summary of daytime (06:00-18:00) average budgets of RO_x radicals (in ppbv h^{-1}) in S1. Primary**
 363 **RO_x sources and sinks are in red and blue. Values in the brackets represent the difference between S1**
 364 **and S0.**

365 3.6.2 Impact on O_3 formation

366 Isoprene is an important precursor of O_3 . To investigate the detailed impact of isoprene
 367 on O_3 formation, the production and loss pathways of O_3 in S1 was also calculated (see
 368 Figure 7 (B)). The simulated daily average level of O_3 dropped to 39.31 ppbv in S1, which is
 369 ~36% lower than that in S0. In addition, the maximum O_3 concentration falls to 80.06 ppbv
 370 in S1, which is ~45.95 ppbv lower than that in S0. Comparisons of S1 and S0 show that the



371 absence of isoprene can reduce all the production and sink pathways of O_3 . For example, the
372 rate of the two major production pathways of O_3 (HO_2+NO and RO_2+NO) decreased by ~21%
373 and ~30%, respectively. This can be attributed to the drop in the concentration of HO_2 and
374 RO_2 radical in S1. As for the depletion of O_3 , the absence of isoprene caused a decrease of
375 0.39 ppbv h^{-1} in the reaction rate of $\text{alkene}+O_3/NO_3$, followed by RO_2+NO_2 (0.16 ppbv h^{-1})
376 and NO_2+OH (0.15 ppbv h^{-1}). Apparently, the absence of isoprene will reduce the total
377 concentrations of alkenes and can further leads to the decrease of RO_2 and OH level, which
378 ultimately slows down the depletion pathways of O_3 . The absence of isoprene also caused a
379 decrease of 5.18 ppbv h^{-1} of the daytime mean net production rate of O_3 . Hence, isoprene
380 chemistry plays an important role in the local O_3 formation at DSL site.

381 3.6.3 Impact on HCHO formation

382 The analysis of S0 revealed the important role of isoprene, aromatics, and alkenes in the
383 production of HCHO. To investigate the chain effect of isoprene chemistry on HCHO
384 production, the major reactions that dominate the formation and depletion of HCHO in S1
385 were also analyzed by OBM model (see Figure 8 (B)). Comparison of S0 and S1 shows that
386 the daily average HCHO decreased by 0.86 ppbv (~15%) when cutting away isoprene
387 chemistry. It is obviously that the drop in HCHO concentration cannot be solely illustrated by
388 the absence of RO (from isoprene). As aforementioned, the absence of isoprene slows down
389 the recycling of RO_x and can further lead to decrease in RO_x concentration. According to the
390 result of OBM analysis, the concentration of CH_3O , RO (from aromatics), RO (from alkanes),
391 and RO (from alkenes) decreased by $24.92 \times 10^1 \text{ molecule cm}^{-3}$, $2.11 \times 10^5 \text{ molecule cm}^{-3}$,
392 $5.12 \times 10^1 \text{ molecule cm}^{-3}$, and $3.94 \text{ molecule cm}^{-3}$, respectively. The drop in the HCHO
393 precursor concentrations ultimately lead to decrease in the daytime reaction rate of $CH_3O +$
394 O_2 , RO (from alkenes) + O_2 , and RO (from aromatics) + O_2 decreased by 0.28 ppbv h^{-1}
395 (~33%), 0.03 ppbv h^{-1} (~16%), and 0.02 ppbv h^{-1} (~8%), respectively. The total daytime
396 formation rate of HCHO dropped to 1.64 ppbv h^{-1} , which was 1.94 ppbv h^{-1} (~54%) lower



397 than that in S0. As a result of the lower HCHO and OH concentration in S1, the daily mean
398 depletion rate of HCHO decreased by 1.79 ppbv h^{-1} (~60%). The absence of isoprene pulls
399 down the daily average HCHO level by only 0.86 ppbv (~15%) because the formation and
400 depletion rates of HCHO decrease at similar rates.

401 **4. Conclusions**

402 Our observations at a rural site of the YRD region from April to June in 2018 captured 5
403 typical local O_3 formation episodes. The detailed atmospheric photochemistry during these
404 episodes were analyzed. Under stagnant condition, the photolysis of OVOCs served as the
405 predominant primary RO_x sources. RO_x achieves efficient recycling with the participation of
406 NO_x . Influenced by the fast RO_x recycling, local O_3 was efficiently produced and
407 accumulated under stagnant conditions. The reactions of RO radicals with O_2 dominate the
408 photochemical formation of HCHO. The higher atmospheric oxidative capacity lead to fast
409 degradation of VOCs, which can further lead to high levels of HCHO at the DSL site.
410 Specifically, the degradation of RO radicals (e.g. ISOP34O, ISOPDO, ISOPAO and ISOPBO)
411 from isoprene oxidation play an important role in the photochemical production of HCHO.
412 To investigate the role of isoprene in RO_x recycle and the formation of secondary pollutant, a
413 sensitivity scenario without isoprene (S1) input was simulated by OBM model. By comparing
414 S1 to the standard simulation, S0, we find that isoprene chemistry is important to local RO_x
415 recycling. The absence of isoprene can obviously decrease the concentrations of OVOCs and
416 the reaction rates in RO_x propagations, and further reduce the concentrations of radicals (e.g.
417 OH, HO_2 , RO_2). Our results indicate that the isoprene chemistry can strongly influence the
418 formation of O_3 and HCHO with the present of NO_x . Removing isoprene can slow down the
419 reaction of HO_2+NO and RO_2+NO by ~21% and ~30%, respectively, and eventually cause
420 ~36% decrease of O_3 . As a result of lower O_3 concentration, average concentration of NO_3
421 dropped by 11% in S1. The absence of isoprene can lead to decrease of RO (from isoprene)



422 and RO_x concentration and cause an obvious drop of HCHO formation (~54%). On the other
423 hand, the decrease in RO_x and HCHO also leads to slower depletion rate of HCHO. Therefore,
424 cutting isoprene can only lead to limited decrease in HCHO concentration. Overall, this study
425 underlines the significant role of isoprene chemistry in radical chemistry, photochemical
426 reactions, and secondary pollutant formation in the atmosphere of the YRD region and
427 provides insights into secondary pollution and its formation mechanisms.

428

429 *Data availability.* The data that support the results are available from the corresponding
430 author upon request.

431

432 *Authorship contribution.* Kun Zhang: Formal analysis, Writing-original draft. Li Li: Writing-
433 review & editing. Ling Huang: Formal analysis. Juntao Huo: Formal analysis. Yusen Duan:
434 Formal analysis. Yuhang Wang: Formal analysis. Yangjun Wang: Formal analysis. Qingyan
435 Fu: Formal analysis.

436

437 *Competing interest.* The authors declare that they have no known competing financial
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440

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