

1 **Explicit modelling of isoprene chemical processing in polluted air**
2 **masses in suburban areas of the Yangtze River Delta region:**
3 **radical cycling and formation of ozone and formaldehyde**
4

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13

14 **Abstract**

15 Ozone pollution has become one of the most severe environmental problems in China in
16 recent years. Our online observations showed that high levels of O₃ were frequently observed
17 in suburban areas of the Yangtze River Delta (YRD) region even there was no obvious ozone
18 transport from the urban regions. To better understand the formation mechanism of local O₃
19 pollution and investigate the potential role of isoprene chemistry in the budgets of RO_x
20 (OH+HO₂+RO₂) radicals, synchronous observations of volatile organic compounds (VOCs),
21 formaldehyde (HCHO) and meteorological parameters were conducted at a suburban site of
22 the YRD region in 2018. Five episodes with elevated O₃ concentrations under stagnant
23 meteorological conditions were identified; an observation-based model (OBM) with the Master
24 Chemical Mechanism was applied to analyze the photochemical processes in these high-O₃
25 episodes. High levels of O₃, nitrogen oxides (NO_x), and VOCs facilitated strong production

26 and recycling of RO_x radicals with the photolysis of oxygenated VOCs (OVOCs) being the
27 primary source. Our results suggest that local biogenic isoprene is important to suburban
28 photochemical processes. Removing isoprene could drastically slow down the efficiency of
29 RO_x recycling and reduce the concentrations of RO_x. The absence of isoprene chemistry could
30 further lead to decrease in the daily average concentration of O₃ and HCHO by 34% and 36%,
31 respectively. This study underlines that the isoprene chemistry in suburban atmosphere
32 becomes important with the participation of anthropogenic NO_x and also provides insights into
33 the radical chemistry that essentially drives the formation of secondary pollutants (e.g. O₃ and
34 HCHO) in suburban YRD region.

35 **Keywords:** Isoprene; Observation-based model (OBM); Radical; Ozone; Yangtze River Delta

36 **1. Introduction**

37 The hydroxyl radical (OH), hydro peroxy radical (HO₂) and organic peroxy radical (RO₂),
38 collectively known as RO_x dominate the oxidative capacity of the atmosphere and hence
39 govern the removal of primary contaminants (e.g. volatile organic compounds (VOCs)) and
40 the formation of secondary pollutants (e.g. ozone (O₃), secondary organic aerosols (SOAs))
41 (Liu et al., 2012;Xue et al., 2016). RO_x radicals can undergo efficient recycling (e.g. OH→
42 RO₂→RO→HO₂→OH) and produce O₃ and oxygenated VOCs (OVOCs) (Liu et al., 2012;Tan
43 et al., 2019;Xue et al., 2016). In addition, the photolysis of OVOCs can in turn produce primary
44 RO₂ and HO₂ radicals, and further accelerate the recycling of RO_x (Liu et al., 2012). The
45 reaction rates of different VOCs with RO_x vary significantly (Atkinson and Arey, 2003;
46 Atkinson et al., 2006). For instance, the reaction rate constants for OH with ethane and ethene
47 are 0.248×10^{-12} (cm molecule⁻¹ s⁻¹) and 8.52×10^{-12} (cm molecule⁻¹ s⁻¹), respectively. Among
48 the hundreds thousands of VOC species, isoprene (C₅H₈, 2-methyl-1,3-butadiene) is one of the
49 most active species, and also the most abundant biogenic VOCs (BVOCs) species globally

50 (Wennberg et al., 2018). Isoprene emissions from biogenic sources have been extensively
51 studied over past decades (Gong et al., 2018) and recent works have switched from emissions
52 to the degradation pathways and the impact of isoprene chemistry on regional forest chemistry
53 (Gong et al., 2018; Wolfe et al., 2016a). Previous studies showed that isoprene could be quickly
54 oxidized by atmospheric oxidants (e.g. OH, O₃ or NO₃) (Wolfe et al., 2016a; Gong et al., 2018;
55 Jenkin et al., 2015). Due to the rapid reaction between OH and isoprene ($100 \times 10^{-12} \text{ cm}^3$
56 $\text{molecule}^{-1} \text{ s}^{-1}$ at 298 K), more than 90% of the total daytime isoprene is removed via this
57 reaction (Wennberg et al., 2018). The reaction between OH and isoprene is initiated by the
58 addition of OH and can generate isoprene hydroxyperoxy radicals (ISOPO₂) (Wennberg et al.,
59 2018; D'Ambro et al., 2017; Liu et al., 2013; Jenkin et al., 2015). ISOPO₂ isomers could then
60 interconvert rapidly due to reversible O₂ addition and are finally removed via reactions with
61 HO₂ or NO (Jenkin et al., 2015; Wolfe et al., 2016a). Hence, the degradation process of
62 isoprene is tightly associated with RO_x recycling. According to He et al. (2019), isoprene
63 chemistry could strongly influence the photochemical formation of O₃, with a relative
64 incremental reactivity (RIR) of ~0.06%/%. In addition to O₃, HCHO is formed via several
65 pathways during the depletion of isoprene (Jenkin et al., 2015; Wolfe et al., 2016a) and HCHO
66 formation is found to be highly sensitive to isoprene (Zeng et al., 2019).

67 The Yangtze River Delta (YRD) region is one of the most developed city-clusters in
68 eastern China and has been suffering from serious O₃ pollution (Zhang et al., 2019; Zhang et
69 al., 2020a; Chan et al., 2017). At the suburban area of YRD, high levels of O₃ were frequently
70 observed (Zhang et al., 2019; Zhang et al., 2020a). A number of studies have been conducted
71 to investigate the relationships between O₃ precursors and O₃ (Chan et al., 2017; Lin et al.,
72 2020; Zhang et al., 2020a; Zhang et al., 2020b), but few have attempted to address the
73 atmospheric oxidizing capacity and radical chemistry involved in these complicated
74 photochemical processes (Tan et al., 2019; Zhu et al., 2020b). Previous studies have pointed

75 out that high levels of O₃ at suburban areas of Shanghai could be attributed to the transport of
76 O₃ or O₃ precursors from urban areas (Lin et al., 2020; Zhang et al., 2019). However, high O₃
77 concentrations were frequently observed in suburban areas under stable meteorological
78 conditions. Given the high vegetation coverage in suburban YRD and weak transport of air
79 masses, the importance of local isoprene chemistry to ozone formation remains unclear.

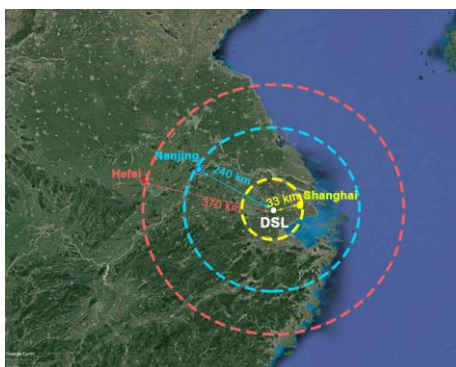
80 In this study, we conducted a comprehensive set of in-situ observations of isoprene,
81 meteorological parameters, and trace gases to understand the important impact of isoprene
82 chemistry on atmospheric photochemical processes in suburban YRD region. An observation-
83 based model (OBM) was used to explore the role of local isoprene chemistry in radical budgets
84 and the formation of O₃ and HCHO. Results from this study can provide insights into the
85 isoprene chemistry in the suburban region of a fast-developing city-cluster.

86 **2. Methodology**

87 **2.1 Measurement site and techniques**

88 The observations were conducted at a supersite (120.98°E, 31.09°N) in the suburban areas
89 of the YRD region (Figure 1). It is located in the west of Shanghai and is close to the Dianshan
90 Lake Scenic area, which has high vegetation coverage. To investigate the local isoprene
91 chemistry and its influence on O₃ and HCHO formation, continuous measurements were
92 conducted from Apr. 7 to Sep. 25, 2018, when the photochemical reactions are active and ozone
93 formation is significant.

94



95

96 **Figure 1. Location of the Dianshan lake supersite (white dot) in the suburban areas of YRD region. This**
 97 **picture was created with Google Earth© on 23rd July 2020.**

98

Table 1. Measurements performed during the ozone season

Species/Parameter	Experimental Technique	Time resolution	Lower Detectable limit
O ₃	Model 49i, Thermo Fischer Scientific, USA	60 s	0.5 ppbv
NO and NO ₂	Model 42i, Thermo Fischer Scientific, USA	60 s	0.4 ppbv
CO	Model 48i, Thermo Fischer Scientific, USA	60 s	40 ppbv
HCHO	AL4021, Aero-Laser, GER	90 s	0.1 ppbv
VOCs species	GC866, Agilent., USA	1 hour	-
Temperature, relative humidity, wind speed and wind direction	Meteorological station, Vaisala, NLD	60 s	-

99

100 The measuring instruments are shown in Table 1. Wind speed (WS), wind direction (WD),
 101 temperature (T), and relative humidity (RH) were simultaneously observed by a meteorological
 102 station (Vaisala., FIN). According to China's air quality standard, several criteria air pollutants
 103 were measured during this experiment. O₃ was measured by an ultraviolet photometric analyzer
 104 (Model 49i, Thermo Fischer Scientific., USA), which has a detection limit of 0.5 ppbv at 60
 105 second resolution. 1 min resolution of nitrogen oxides (NO and NO₂) data were simultaneously
 106 observed by a chemiluminescence instrument (Model 42i, Thermo Fischer Scientific., USA),
 107 which has a detection limit of 0.4 ppbv. Carbon monoxide was monitored by a gas filter
 108 correlation infrared absorption analyzer (Model 48i, Thermo Fischer Scientific., USA), which
 109 has a detection limit of 0.04 ppm. All the online instruments used for gas analyzer were auto-
 110 zero every day, and were multi-point calibrated every month. All the instruments used for the

111 online observation were housed on top of a 5-floor-high building, which was about 15 m above
112 the ground level.

113 A total of 55 VOC species, including 28 alkanes, 10 alkenes, 16 aromatics and acetylene
114 were continuously analyzed at our sampling site by two online gas chromatograph with flame
115 ionization detector (GC-FID) systems (GC-866 airmoVOC C₂-C₆ #58850712 and airmoVOC
116 C₆-C₁₂ #283607112, Agilent., USA) with a time resolution of 1 hour during our experiment.
117 Ambient samples are directly inhaled into this system by a pump. Low carbon VOCs (C₂-C₆)
118 are captured by a low temperature (-10 °C) preconcentration system, while high carbon VOCs
119 are concentrated by a built-in room temperature preconcentration system. Then the
120 preconcentration systems are heated and desorb VOCs, which are then carried into
121 chromatographic columns by helium. Individual VOCs separated in the columns are eventually
122 detected by FID systems. Formaldehyde (HCHO) was continuously measured by a Hantzsch
123 fluorescence technique (AL4201, Aerolaser GmbH., GER), which is based on fluorometric
124 Hantzsch reaction in the liquid phase, requiring the quantitative transfer of HCHO from gas
125 phase to liquid phase. A Hantzsch reagent (acetylacetone) was used in this instrument.

126 **2.2 Observation-based model**

127 A user-friendly zero-dimensional (0-D) box model (F0AM) was used to simulate the
128 chemical processes in the atmosphere in this study. This model was developed by Wolfe et
129 al.(2016b) based on University of Washington Chemical Model (UWCM). Dry deposition,
130 aloft exchange, and atmospheric dilution were considered in this model. We chose the Master
131 Chemical Mechanism (MCM) v3.3.1 as the chemical mechanism with more than 5,800
132 chemical species and 17,000 reactions, which enables a detailed description of the complex
133 reactions. In addition to gas-phase reactions, several heterogenous processes including the
134 uptake of HO₂, N₂O₅ and HCHO on aerosol surface, and heterogenous source of HONO were
135 also considered in our simulation. These reactions rate constants and uptake coefficient were

136 obtained from the study of Riedel et al. (2014), Xue et al. (2014) and Li et al. (2014). Since
 137 key parameters such as aerosol surface areas (S_A) and particle diameters (r) were not measure
 138 during our observation, an average S_A ($640 \text{ nm}^2/\text{cm}^3$) was obtained from the field campaign in
 139 Shanghai (Wang et al., (2014)).

140 Table 2. **Heterogenous reactions and associated rate constants used in the OBM model**

Reactions	Reaction rate constant	Reference
$N_2O_5 \rightarrow CLNO_2 + HNO_3$	$\gamma\omega S_A/4$ (for $CLNO_2$ formation)	Riedel et
	$(2 - \phi)\gamma\omega S_A/4$ (for HNO_3 formation)	al. (2014)
$NO_2 \rightarrow HONO$	$k_g = \frac{1}{8} \times \omega\gamma_g \left(\frac{S}{V}\right)$	Xue et al.
	$k_a = \frac{1}{4} \omega\gamma_a S_A$	(2014)
$HO_2 \rightarrow products$	$k = \left(\frac{r}{D_g} + \gamma\right)^{-1} S_A$	Xue et al. (2014)
$HCHO \rightarrow products_1$	$k = \frac{1}{4} \omega\gamma S_A$	Li et al. (2014)

γ = uptake coefficient for the given reactant with aerosol surface area; ϕ = product yield; ω =mean molecular speed of the given reactant (m/s); S_A =RH corrected aerosol surface area concentration (nm^2/cm^3); r =surface-weighted particle radius.

141

142 Photolysis frequencies (J values) were calculated by a trigonometric parameterization
 143 based on solar zenith angle (SZA):

$$J = I \cos(SZA)^m \exp(-n \sec(SZA)) \quad (1)$$

144 where I , m and n are constants unique to each photolysis reaction, derived from least-squares
 145 fits to J values computed with fixed solar spectra and literature cross-section and quantum
 146 yields (Wolfe et al., 2016b). Hourly averaged concentrations of speciated VOCs (except
 147 HCHO), NO, NO₂, CO and meteorological parameters (such as T, RH and P) were used to
 148 constrain the F0AM model. Nitrous acid (HONO) was not measured during our observation.

149 Therefore, it was fixed as 2% of the observed NO₂ concentration. This constant ratio is well
150 observed in different field studies and performed well in previous box model studies (Tan et
151 al., 2019). Before each simulation, the model will run 3 days as spin up to reach a steady state
152 for unmeasured species (e.g., OH and NO₃ radicals). The comparison of simulated and
153 observed O₃ and HCHO concentrations is shown in Figure S1 and Figure S2. The index of
154 agreement (IOA), mean bias (MB) and normalized mean bias (NMB) are frequently used to
155 estimate the model performance. These three parameters can be calculated by Equation (2) to
156 (4), where S_i, O_i, and \bar{O} are the simulated, observed, and average observed value of the target
157 compound. In this study, the IOA, MB and NMB of O₃ was 0.90, 0.76 and 10%, respectively.
158 This result suggests that the model can reasonably reproduce the variations of O₃ and could be
159 used for further analysis. As for HCHO, the IOA, MB, and NMB was 0.74, 2.43 and 48%,
160 respectively. In general, the model overestimated HCHO concentration, especially on July 29
161 and July 30. According to previous studies, the inconsistency between simulated and observed
162 HCHO could be caused by the uncertainties in the treatment of dry deposition, faster vertical
163 transport, uptake of HCHO, and fresh emissions of VOCs precursors (Li et al., 2014). But the
164 result still provides valuable information of secondary formation of HCHO at suburban area.
165 To quantify the atmospheric oxidative capacity (AOC) changes in response to isoprene
166 chemistry, two parallel scenarios (S0 and S1) were conducted with isoprene chemistry disabled
167 in S1. In both cases, identical chemical mechanism and meteorological conditions were used
168 to drive model simulations. Through a comparative analysis of the scenarios, the impact of
169 isoprene chemistry on AOC and secondary formation of O₃ and HCHO can be obtained.

$$IOA = 1 - \frac{\sum(S_i - O_i)^2}{\sum(|S_i - \bar{O}| + |O_i - \bar{O}|)^2} \quad (2)$$

$$MB = \frac{\sum(S_i - O_i)}{N} \quad (3)$$

$$NMB = \frac{\sum(S_i - O_i)}{\sum O_i} \times 100\% \quad (4)$$

170 **3. Results and discussions**

171 **3.1 Overview of the observations**

172 To investigate the impact of local chemistry on ozone formation, five days with low daily
173 average wind speed (<2m/s) and high maximum daily 8-h average (MDA8) O₃ concentration
174 (>74.7 ppb) were identified as typical local chemistry cases. Figure 2 shows the time series of
175 observed meteorological parameters (P, T, and RH), trace gases (NO, NO₂ and O₃), isoprene
176 and HCHO on selected days. During those episodes, the air masses reaching the site were
177 mainly from southeast and southwest (Figure 2). The weak wind was not conducive to the
178 regional transportation of air pollutants. The observed O₃, NO₂, NO, CO, and TVOC ranged
179 from 1.40 to 155.40 ppbv (52.72 ± 44.43 ppbv, average value, the same below), 5.36 to 57.95
180 ppbv (21.58 ± 12.88 ppbv), 0.75 to 54.51 ppbv (5.40 ± 8.13 ppbv), 400 to 960 ppbv (597 ± 153
181 ppbv), and 2.34 to 20.33 ppbv (7.28 ± 4.32 ppbv) respectively. During the five episodes, the
182 average concentrations of alkanes, alkenes, and aromatics were 13.97 ± 9.12 , 3.27 ± 2.31 , and
183 4.93 ± 2.69 ppbv, which were about 53%, 18%, and 50% higher than of the whole observation.
184 The conditional probability function (CPF) is applied to exhibit the relationship between high
185 O₃ concentrations and wind (Figure 3). The detailed description of CPF can be found in
186 supplemental information (Text S1). The result suggests that high O₃ concentrations (>131 ppb)
187 was usually observed when the site was influenced by weak south wind. This implies that high
188 O₃ was most likely formed locally. Although this site is far away from urban areas, high levels
189 of NO were found during early morning, which is likely caused by nearby fresh emissions. As
190 for NO₂, only one peak was found at dusk. This was different from the results in urban areas
191 (Zhang et al., 2019). It is worth noting that NO₂ and O₃ concentrations were high even during
192 nighttime, suggesting that the AOC remained high at nighttime. The daily average isoprene
193 concentrations were 0.37 ± 0.36 ppbv, which is comparable to that observed by Gong et al.

194 (2018) at a forested mountaintop site (0.287 ± 0.032 ppbv). To roughly estimate the influence
195 of isoprene on atmospheric oxidation capability, we adopted the approach given in the study
196 of Zhu et al. (2020) to calculate the OH reactivity (k_{OH}). The result suggested that isoprene,
197 accounting for ~19% of the total k_{OH} , was the most significant VOC specie from the perspective
198 of k_{OH} , with an average value of 0.89 ± 0.44 s⁻¹. This indicates the significant role of isoprene
199 in the photochemistry in suburban area. The average HCHO was 5.01 ± 3.80 ppbv, which was
200 ~2 times of that observed at a rural site of Hong Kong (Yang et al., 2020). It is worth noting
201 that HCHO could reach 18.69 ppbv at midday. Based on the explicit calculation, the total
202 concentration of OVOC was obtained. Due to the complexity of OVOC formation, which could
203 have hundreds of precursors for just one OVOC specie, and the complicated chain reactions
204 converting VOCs to OVOCs, it is difficult to give the accurate relationship between VOCs to
205 OVOCs. Since VOCs were mainly oxidized by OH and O₃ during daytime, in this study, we
206 chose multi-linear regression model (given in Eq.(5)) to explore the roughly relationship
207 between VOCs and simulated OVOCs.

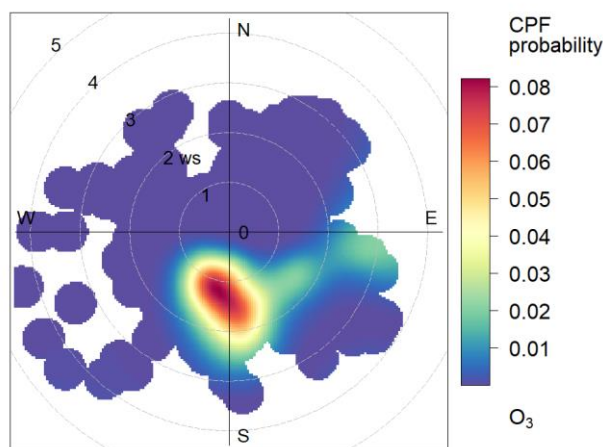
$$[OVOC] = \beta_0 + \beta_1[Alkane] + \beta_2[Alkene] + \beta_3[Aromatic] + \beta_4[OH] + \beta_5[O_3] \quad (5)$$

208 where β_0 , β_1 , β_2 , β_3 , β_4 , and β_5 are the coefficient from linear regression; [OVOC] and [OH] are
209 the simulated concentration of OVOC and OH, respectively; [Alkane], [Alkene], [Aromatic],
210 [O₃] are the observed concentration of alkanes, alkenes, aromatics, and O₃, respectively. The
211 Sig value and statistical reliability criteria (R) was 0.000 and 0.853 (shown in Table S2),
212 respectively, indicating that the linear relationship represented by equations (5) is statistically
213 reliable. The β_1 , β_2 , β_3 was 0.027, 0.623, and 0.820, respectively, suggesting that alkenes and
214 aromatics are significant for the simulated OVOC concentration.



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Figure 2. Time series of hourly averages for trace gases, isoprene, HCHO, and meteorological parameters.



CPF at the 95th percentile (=131)

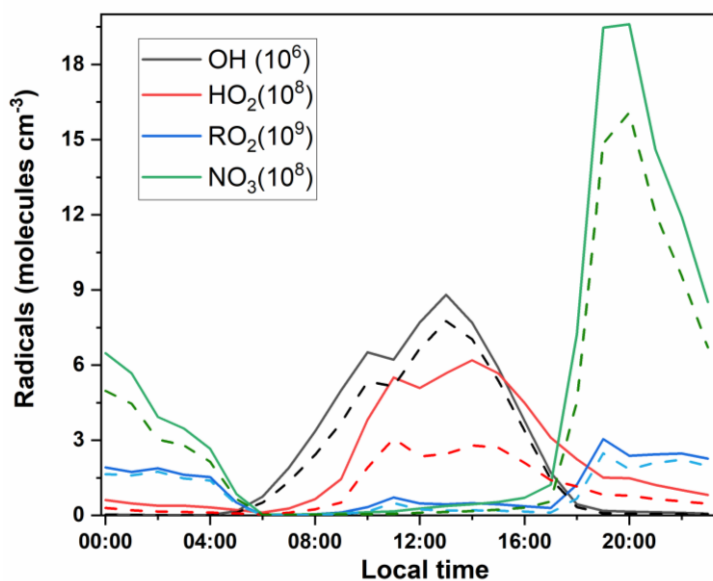
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Figure 3. CPF polar plot of O₃ at DSL station.

220 3.2 Simulated concentrations of radicals

221 Figure 4 shows the simulated average diurnal variation of major radicals in the base
222 scenario (S0). The average concentrations of OH, HO₂, RO₂, and NO₃ were estimated at 2.11
223 $\times 10^6$, 1.04×10^8 , 0.90×10^9 , and 3.49×10^8 molecules cm⁻³, respectively. The simulated daily

224 average OH concentration lies between the simulated values during the summer in Beijing (9
225 $\times 10^6$ molecules cm^{-3}) and the simulated value at a suburban site in Hong Kong in 2013 ($1.5 \pm$
226 0.2×10^6 molecules cm^{-3}) (Liu et al., 2019; Xue et al., 2016). In addition, the average simulated
227 daytime OH concentration was $\sim 50\%$ lower than that simulated at a forested mountaintop site
228 in southern China (Gong et al., 2018). To verify the performance of OBM model, regional
229 mixing ratios of OH during daytime were also calculated by a parameterization method using
230 measured ethylbenzene and *m,p*-xylene ratios (see Text S2). The calculated average regional
231 concentrations of OH ($8.39 \pm 5.11 \times 10^6$ molecules cm^{-3}) was in the same magnitude of the
232 OBM-simulated result ($4.59 \pm 5.11 \times 10^6$ molecules cm^{-3}), suggesting that the OBM-simulated
233 radical concentration is reliable. The maximum HO₂ concentration simulated for DSL site (6.19
234 $\times 10^8$ molecules cm^{-3}) was close to that reported in Beijing (6.8×10^8 molecules cm^{-3}) (Liu et
235 al., 2012), but was $\sim 32\%$ higher than that in Wuhan (4.7×10^8 molecules cm^{-3}) (Zhu et al.,
236 2020a). Pretty high levels of simulated NO₃ (as high as $\sim 19 \times 10^8$ molecules cm^{-3}) was found
237 during nighttime. The average simulated nocturnal NO₃ concentration was 8.80×10^8 molecule
238 cm^{-3} , which was $\sim 47\%$ higher than that simulated in the study of Gong et al.(2018). As
239 aforementioned, during nighttime, pretty high levels of NO₂ (27.71 ppbv) and O₃ (30.05 ppbv)
240 was observed, which favored the formation of NO₃. Interestingly, a high level of RO₂ was also
241 found during nighttime. This result is different from the study of Liu et al. (2012), which found
242 the maximum value of RO₂ during daytime. By separating the formation pathways of RO₂, we
243 found that during nighttime, over 70% RO₂ was produced via the oxidation of VOCs by NO₃
244 radical, suggesting that the nighttime chemistry in the suburban site was also very important.
245



246
 247 **Figure 4. Simulated average diurnal variation of OH, HO₂, RO₂ and NO₃ in S0 (solid lines) and S1 (dash**
 248 **lines).**

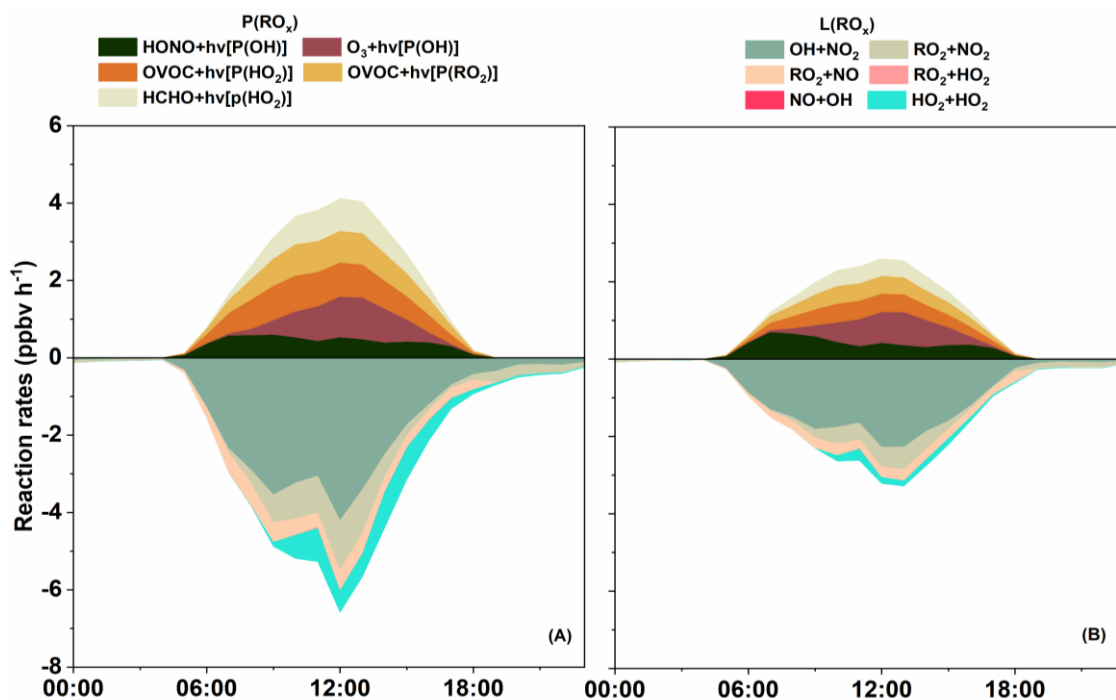
249 **3.3 Recycling of RO_x radicals**

250 Figure 5(A) shows the primary sources of RO_x in S0 and the detailed daytime budget of
 251 RO_x. Minor RO_x sources, e.g. ozonolysis of alkenes, are not shown. The photolysis of O₃ was
 252 the predominant primary source of OH, with a daytime mean production rate of 0.50 ppbv h⁻¹,
 253 which was comparable to that found by Liu et al. (2012) in Beijing, but was 0.40 ppbv h⁻¹ lower
 254 than the result in the study of Xue et al. (2016). Another important OH source is the photolysis
 255 of HONO, contributing 0.32 ppbv h⁻¹ of daytime OH production in our simulation. This value
 256 is much lower than the results of Liu et al. (2019) and Xue et al. (2016). Such low value was
 257 most likely caused by the excessive constrain on HONO since HONO was not directly
 258 monitored during our experiment. Sensitive studies were conducted to quantify the influence
 259 different of HONO/NO₂ ratio on radical recycling (Text S3, Figure. S1 and Table S1). As
 260 expected, a lower HONO/NO₂ ratio leads to a lower HONO concentration, and subsequent less
 261 OH reduce generated from the photolysis of HONO. The sensitive studies show that when
 262 HONO/NO₂ ratio is 0.005, the daytime OH level could decrease by 15.28%. Vice versa, a
 263 higher HONO/NO₂ (e.g., 0.04) can promote OH concentration by 14.08%. This result indicates

264 that the photolysis of HONO is essential to the generation of OH, and therefore a simultaneous
265 measurement of HONO is highly recommended for the analysis of local radical recycling in
266 the future. As for HO₂, the photolysis of OVOC (excluding HCHO) is the predominant source
267 with a daytime mean production rate of 0.65 ppbv h⁻¹ and maximum reaching 0.92 ppbv h⁻¹,
268 which is comparable to Xue et al. (2016). The photolysis of HCHO can also contribute 0.48
269 ppbv h⁻¹ to the daytime production of HO₂, which is close to the results of Xue et al. (2016).
270 As for RO₂, the photolysis of OVOC was the largest source (0.57 ppbv h⁻¹), which was
271 relatively lower than the results found at urban site (Liu et al., 2012). From the RO_x perspective,
272 the daytime primary radical production in DSL site was dominated by the photolysis of OVOC
273 (except for HCHO), followed by the photolysis of HCHO and O₃. But the photolysis of HONO
274 can become the overriding RO_x source around run rising, which suggests that HONO can be
275 an important OH reservoir species during nighttime. Summing up all the sources of RO_x gives
276 a total primary daytime RO_x production rate of 2.55 ppbv h⁻¹ (0.84 ppbv h⁻¹ for OH, 1.14 ppbv
277 h⁻¹ for HO₂, and 0.57 ppbv h⁻¹ for RO₂), which was 61~69% lower than those in Beijing (6.6
278 ppbv h⁻¹, Liu et al. (2012)) and Hong Kong (8.11 ppbv h⁻¹, Xue et al. (2016)), indicating that
279 the recycling of RO_x in Beijing and Hong Kong could be much reactive.

280 RO_x radicals are ultimate removed from the atmosphere via deposition of radical reservoir
281 species, e.g. H₂O₂, HNO₃, and ROOH (Liu et al., 2012). The terminate processes of RO_x was
282 dominated by their reactions with NO_x. Specifically, the reaction of OH+NO₂, RO₂+NO₂,
283 RO₂+NO, forming HNO₃, RO₂NO₂, and RONO₂, accounting for 2.42, 0.56, and 0.41 ppbv h⁻¹
284 of the RO_x radical sink during daytime. This is consistent with the understanding that reactions
285 with NO_x usually dominate the radical sink in high NO_x environments (Xue et al., 2016; Liu
286 et al., 2012). In addition, RONO₂ and RO₂NO₂ could in turn react with OH, leading to 0.41
287 ppbv h⁻¹ of daytime OH sinks (Figure 6). Summing up the primary sources and sinks gives a

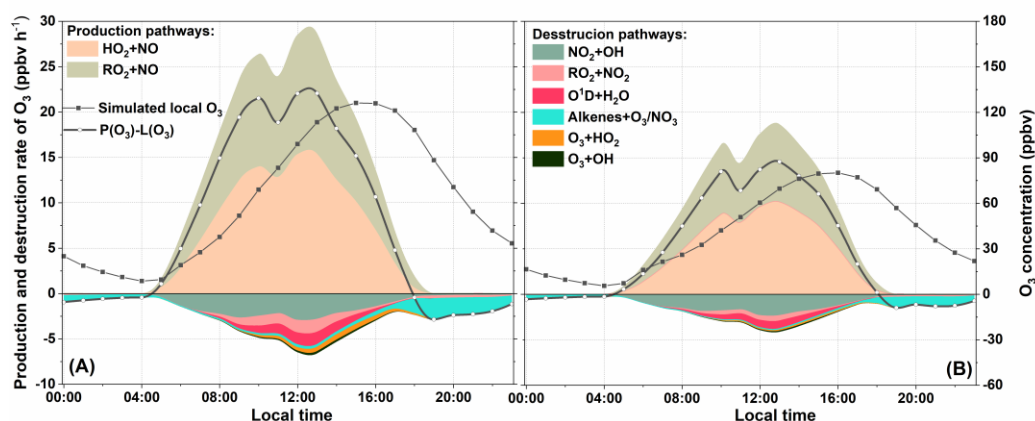
288 negative value of net RO_x production, suggesting that the RO_x was in a stage of gradual
 289 depletion.



290
 291 **Figure 5. Simulated primary daytime sources and sink of RO_x in S0 (A) and S1 (B).**

292 The daytime (6:00-18:00) average budget of RO_x is shown in Figure 6, with primary
 293 sources of RO_x in red, sinks of RO_x in blue, and recycling processes in black. In the recycling
 294 of RO_x, the production of OH was dominated by the reaction of HO₂+NO (8.29 ppbv h⁻¹). As
 295 for RO₂, it was produced by the reaction of OH with OVOC (3.02 ppbv h⁻¹), alkyl (RH) (1.21
 296 ppbv h⁻¹), and peroxides (0.14 ppbv h⁻¹). The reaction of RO₂+NO can result in strong
 297 production of RO (3.87 ppbv h⁻¹). The reaction of RO and O₂ was the major contributor to
 298 HO₂ production, followed by the reaction of OH with CO (1.89 ppbv h⁻¹), OVOC (1.59 ppbv
 299 h⁻¹), and RH (0.15 ppbv h⁻¹). It should be noted that the top two fast reactions within the
 300 recycling of RO_x (HO₂+NO and RO₂+NO) were related to NO_x. As mentioned in the study
 301 of Liu et al. (2012), this result could be mainly due to the abundance of NO (e.g. ~50 ppbv in
 302 the morning). Obviously, these recycling processes dominate the total production of OH, HO₂
 303 and RO₂ radicals. As suggested in the study of Xue et al. (2016) and Liu et al. (2012), the

320 and NO₂, the sink of O₃ was due to several reactions leading to the destruction of O₃ and NO₂.
 321 In our cases, the reaction of NO₂+OH becomes the predominant scavenging pathways of O₃,
 322 with an average daytime reaction rate of 1.89 ppbv h⁻¹ (49%, percentage of the total O₃ sink
 323 rate, same below). This is comparable to the study of Liu et al. (2012 and 2019). The reaction
 324 of RO₂+NO₂ was the second contributor to O₃ sink, with a mean contribution of 0.62 ppbv h⁻¹
 325 (16%). Other pathways, e.g. photolysis of O₃, ozonolysis of alkenes, and O₃+HO₂, together
 326 contributed 1.11 ppbv h⁻¹ of the total sink rate of O₃ during daytime. The daytime mean L(O₃)
 327 was 3.87 ppbv h⁻¹, which was ~22% of P(O₃), suggesting that O₃ could efficiently accumulate
 328 during daytime. The net production of O₃ (P(O₃)-L(O₃)) is also shown in Figure 7. The
 329 maximum O₃ concentration was found at around 16:00, which was also observed in other
 330 suburban sites (Zong et al., 2018; Zhang et al., 2019). It is worth noting that, the reaction of
 331 alkenes+O₃/NO₃ serves as an important pathway of O₃ sink during nighttime (as high as 2.30
 332 ppbv h⁻¹).



333
 334 **Figure 7. Simulated average diurnal profiles of O₃ formation and sink rates (ppbv h⁻¹) in S0 (A) and S1**
 335 **(B).**

336 **3.5 Formation and sink of HCHO**

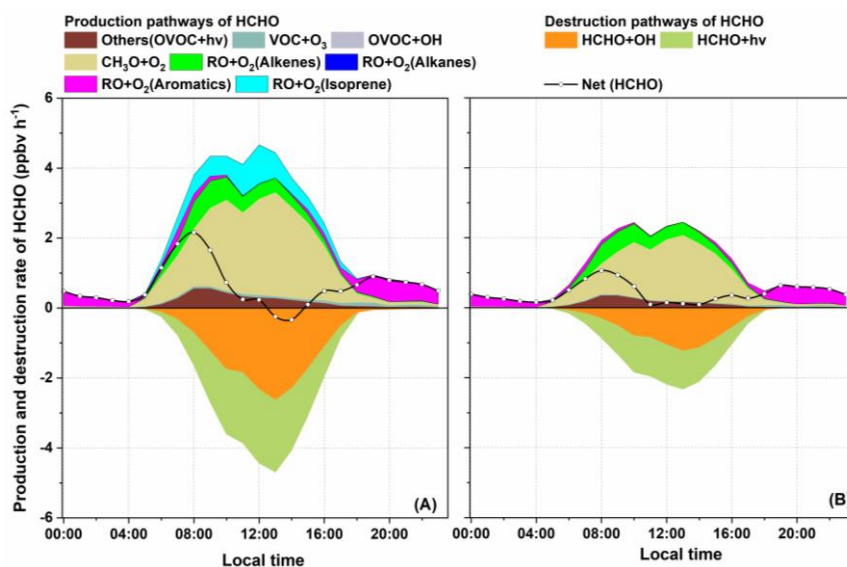
337 As aforementioned, high levels of HCHO was observed at DSL. Figure 8 (A) shows the
 338 production and sink pathways of HCHO in S0. On average, HCHO formation was dominated
 339 by the reaction of RO+O₂, accounting for ~90% of the total production rate. Further

340 classification of RO+O₂ pathway suggested that the oxidation of CH₃O made a significant
341 contribution of ~47%, followed by RO (from isoprene) + O₂ reaction (12%) and RO (from
342 aromatics) + O₂ reaction (~11%). This result is comparable to the study of (Yang et al., 2020;
343 Yang et al., 2018). It is notable that the reaction of RO (from aromatics) + O₂ could become
344 the predominant pathway of HCHO production during nighttime. This could be attributed to
345 the high level of NO₃ during nighttime, by which styrene could be quickly oxidized and
346 generate N-containing RO radicals, and furtherly generate HCHO. During daytime, isoprene
347 became the most important VOC specie of HCHO production, with a mean rate of 0.48 ppbv
348 h⁻¹. As mentioned, this site is surrounded by highly vegetated areas, which can provide
349 abundant biogenic isoprene. During daytime, over 90% of isoprene was oxidized by OH
350 radicals (Figure S4). According to MCMv3.3.1, several RO₂ species (e.g. ISOP34O₂,
351 ISOPDO₂, ISOPCO₂, CISOPAO₂, ISOPAO₂) can be generated during the OH-initiated
352 degradation process of isoprene (Jenkin et al., 2015). With the present of NO, isoprene-
353 originated RO₂ can transfer into RO (e.g. ISOPDO, ISOP34O, ISOPAO). The subsequent
354 degradation processes of isoprene-related RO, especially ISOP34O, ISOPDO, ISOPAO and
355 ISOPBO, are tightly related to the formation of HCHO (Jenkin et al., 2015). Other sources of
356 HCHO, such as the reaction between VOC and O₃, photolysis of OVOC and the reaction of
357 OVOC+OH only contributed minor amount of the total production rate during whole day.

358 As for HCHO depletion, the photolysis of HCHO and the reaction of HCHO+OH was the
359 two dominate pathways, accounting for ~52% and ~48% of the total depletion rate, respectively.
360 The net HCHO production rate (equals to P(HCHO) + L(HCHO)) was also shown in Figure 8.
361 After sunrise, the net production rate of HCHO raised gradually until 8:00, when it reached the
362 maximum rate (1.6 ppbv h⁻¹). This result is comparable to the study of Yang et al. (2018). At
363 around 12:00, the net(HCHO) dropped to ~0 ppbv h⁻¹, that was roughly consistent with our
364 observation, which shows that the HCHO peak occurs at around 12:00. Between 13:00 and

365 14:00, a negative net(HCHO) was found. Although the reaction of RO+O₂ quickly produced
 366 HCHO at afternoon, the depletion pathways, especially the photolysis of HCHO, became more
 367 competitive, leading to the net reduction of HCHO. This also indicated that strong
 368 photochemical reactions do not monotonously profit the accumulation of HCHO, it can also
 369 constrain high HCHO levels in certain situations. After 14:00, the photolysis of HCHO dropped
 370 rapidly and the net depletion of HCHO back to ~0 ppbv h⁻¹ at around 15:00. The daytime net
 371 HCHO production rate was 0.70 ppbv h⁻¹, which was comparable to result of Yang et al. (2018).

372 The above analysis indicates that the photolysis of OVOC, HCHO, O₃ and HONO was
 373 the primary source of RO_x, which offers high oxidizing environment for the degradation of
 374 VOCs. As a typical by-product in the degradation of several VOCs, HCHO can be quickly
 375 formatted during daytime. The insight into detailed photochemical processes shows the
 376 important role of isoprene in the formation of HCHO.



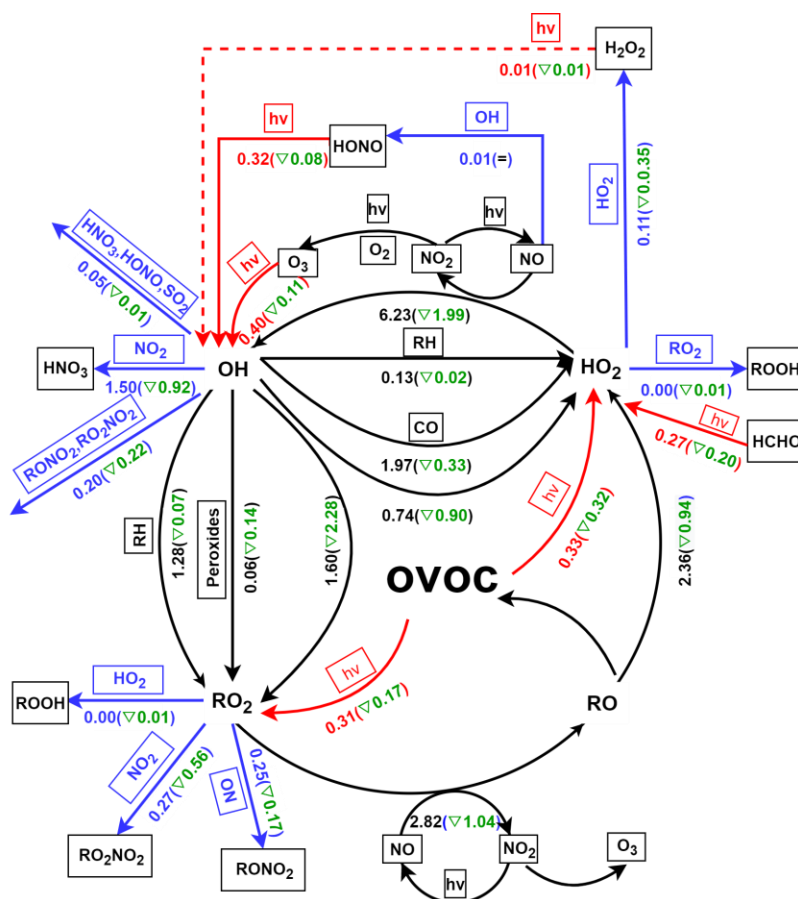
377
 378 **Figure 8. Simulated average diurnal profiles of net rate (net (HCHO)), breakdown HCHO production**
 379 **rate and sink rate (ppbv h⁻¹) in S0 (A) and S1 (B).**

380 3.6 Impacts of isoprene chemistry on photochemistry

381 3.6.1 Impact on RO_x budget

382 As aforementioned, the degradation of isoprene is tightly related to the cycling of RO_x.
 383 To roughly explain the impact of isoprene chemistry on RO_x budget, we carried out a parallel

384 simulation (S1) where isoprene chemistry is disabled (see in Figure 9). The diurnal variation
385 of OH, HO₂, RO₂ and NO₃ in S1 is also shown in Figure 4 (B) which clearly suggests the
386 decline in RO_x and NO₃ without isoprene input. To investigate the underlying causes, we
387 calculated the production rate of RO_x (P(RO_x)) and loss rate of RO_x (L(RO_x)) in S1,
388 respectively (Figure 5 (B)). From the comparison, we found most of the reaction rates in P(RO_x)
389 and L(RO_x) showed a decrease trend in S1, suggesting that the absence of isoprene slows down
390 the RO_x recycling. The photolysis of OVOC (0.67 ppbv h⁻¹) is still the predominant primary
391 source of RO_x. However, without isoprene, the photolysis rate of OVOC decreased by 0.49
392 ppbv h⁻¹. The total production and depletion rate of OH dropped to 6.96 and 7.51 ppbv h⁻¹,
393 respectively. Although the absence of isoprene could reduce the consumption of OH, the OH
394 concentration would be reduced by ~16% compared to S0, suggesting that the amount of OH
395 produced via isoprene chemistry is large enough to compensate for the shift from OH to peroxy
396 radicals in the RO_x family. As for RO₂, the daytime production and sink rate falls to 3.25 and
397 3.34 ppbv h⁻¹, respectively. This means the concentration of RO₂ would be in a stage of gradual
398 decrease. In addition, the absence of isoprene could also reduce RO₂ concentration by ~20%,
399 suggesting that isoprene was an important source of RO₂ at DSL site. As for HO₂, drastic
400 decrease of ~53% was found in S1. The above-mentioned decrease in RO_x obviously could not
401 be explained solely by the remove of isoprene-related radicals. Inspection of the model results
402 shows that OVOC concentrations decreased drastically (~41%) after cutting isoprene (e.g. ~37%
403 decrease in formaldehyde, ~65% decrease in methylglyoxal, ~51% decrease in glyoxal, ~100%
404 decrease in methacrolein (MACR), and ~100% decrease in methyl vinyl ketone (MVK)). The
405 decrease in OVOC can further pull down substantial amount of primary RO₂ and HO₂ (Figure
406 6 and Figure 9). It is interesting to note that, subtracting isoprene also cause drop of NO₃
407 (~23%). This result can be contributed to the decrease of secondary production of O₃ (~35%),
408 which can further reduce the formation of NO₃, especially during nighttime.



409
 410 **Figure 9. Summary of daytime (06:00-18:00) average budgets of RO_x radicals (in ppbv h⁻¹) in S1. Primary**
 411 **RO_x sources and sinks are in red and blue, respectively, and the black lines represent the processes in**
 412 **RO_x and NO_x recycling. Values in the brackets represent the difference between S1 and S0.**

413 3.6.2 Impact on O₃ formation

414 To investigate the detailed impact of isoprene on O₃ formation, the production and sink
 415 pathways of O₃ in S1 was also calculated (see Figure 7 (B)). In S1, the simulated maximum
 416 and daily average level of O₃ dropped to 84.95 and 41.23 ppbv, respectively, which is ~35%
 417 and ~34% lower than that in S0. Comparisons of S1 and S0 show that the absence of isoprene
 418 can reduce all the production and sink pathways of O₃. For example, the rate of the two major
 419 production pathways of O₃ (HO₂+NO and RO₂+NO) decreased by ~37% and ~45%,
 420 respectively. This can be attributed to the drop in the concentration of HO₂ and RO₂ radical in
 421 S1. As for the depletion of O₃, the absence of isoprene caused a decrease of 0.31 ppbv h⁻¹ in
 422 the reaction rate of alkene+O₃/NO₃, followed by RO₂+NO₂ (0.22 ppbv h⁻¹) and NO₂+OH

423 (0.265 ppbv h⁻¹). Apparently, the absence of isoprene will reduce the total concentrations of
424 alkenes and can further leads to the decrease of RO₂ and OH level, which ultimately slows
425 down the depletion pathways of O₃. Eventually, the absence of isoprene caused a decrease of
426 5.78 ppbv h⁻¹ in the daytime mean net production rate of O₃. Hence, isoprene chemistry plays
427 an important role in the local O₃ formation at DSL site.

428 3.6.3 Impact on HCHO formation

429 The analysis of S0 revealed the important role of isoprene, aromatics, and alkenes in the
430 production of HCHO. To investigate the chain effect of isoprene chemistry on HCHO
431 production, the major reactions that dominate the formation and depletion of HCHO in S1 were
432 also analyzed by OBM model (see Figure 8 (B)). Comparison of S0 and S1 shows that the daily
433 average HCHO decreased by 2.90 ppbv (~39%) when cutting away isoprene chemistry. It is
434 obviously that the drop in HCHO concentration cannot be solely illustrated by the absence of
435 RO (from isoprene). As aforementioned, the absence of isoprene slows down the recycling of
436 RO_x and can further lead to decrease in RO_x concentration. According to the result of OBM
437 analysis, the concentration of CH₃O, RO (from aromatics), RO (from alkanes), and RO (from
438 alkenes) decreased by 2.70×10² molecule cm⁻³, 1.59×10⁵ molecule cm⁻³, 3.35×10¹ molecule
439 cm⁻³, and 3.44 molecule cm⁻³, respectively. The drop in the HCHO precursor concentrations
440 ultimately lead to decrease in the daytime reaction rate of CH₃O + O₂, RO (from alkenes) + O₂,
441 and RO (from aromatics) + O₂ decreased by 0.66 ppbv h⁻¹ (~36%), 0.06 ppbv h⁻¹ (~16%), and
442 0.06 ppbv h⁻¹ (~40%), respectively. The total daytime formation rate of HCHO dropped to 1.71
443 ppbv h⁻¹, which was 1.66 ppbv h⁻¹ (~49%) lower than that in S0. As a result of the lower HCHO
444 and OH concentration in S1, the daily mean depletion rate of HCHO decreased by 1.25 ppbv
445 h⁻¹ (~49%). Finally, the absence of isoprene pulls down the daily average HCHO level by
446 1.61ppbv (~36%).

447 **4. Conclusions**

448 Our observations at a suburban site of the YRD region from April to June in 2018 captured
449 5 typical local O₃ formation episodes. The detailed atmospheric photochemistry during these
450 episodes were analyzed. Under stagnant condition, the photolysis of OVOC served as the
451 predominant primary RO_x sources. RO_x achieves efficient recycling with the participation of
452 NO_x. Influenced by the fast RO_x recycling, local O₃ was efficiently produced and accumulated
453 under stagnant conditions. The reactions of RO radicals with O₂ dominate the photochemical
454 formation of HCHO. The higher atmospheric oxidative capacity lead to fast degradation of
455 VOCs, which can further lead to high levels of HCHO at the DSL site. Specifically, the
456 degradation of RO radicals (e.g. ISOP34O, ISOPDO, ISOPAO and ISOPBO) from isoprene
457 oxidation play an important role in the photochemical production of HCHO. To investigate the
458 role of isoprene in RO_x recycle and the formation of secondary pollutant, a sensitivity scenario
459 without isoprene (S1) input was simulated by OBM model. By comparing S1 to the standard
460 simulation (S0), we find that isoprene chemistry is important to local RO_x recycling. The
461 absence of isoprene can obviously decrease the concentrations of OVOC and the reaction rates
462 in RO_x propagations, and further reduce the concentrations of radicals (e.g. OH, HO₂, RO₂).
463 Our results indicate that the isoprene chemistry can strongly influence the formation of O₃ and
464 HCHO with the present of NO_x. Removing isoprene can slow down the reaction of HO₂+NO
465 and RO₂+NO by ~37% and ~45%, respectively, and eventually cause ~34% decrease of O₃.
466 As a result of lower O₃ concentration, average concentration of NO₃ dropped by 23% in S1.
467 The absence of isoprene can lead to decrease of RO (from isoprene) and RO_x concentration
468 and cause an obvious drop of HCHO formation (~49%). Overall, this study underlines the
469 significant role of isoprene chemistry in radical chemistry, photochemical reactions, and
470 secondary pollutant formation in the atmosphere of the YRD region and provides insights into
471 secondary pollution and its formation mechanisms.

472

473 *Data availability.* The data that support the results are available from the corresponding author
474 upon request.

475

476 *Authorship contribution.* Kun Zhang: Formal analysis, Methodology, Writing-original draft.
477 Ling Huang: Writing-review. Qing Li: Formal analysis. Juntao Huo: Formal analysis, Data
478 curation. Yusen Duan: Formal analysis, Data curation. Yuhang Wang: Writing-review.
479 Yangjun Wang: Formal analysis. Qingyan Fu: Formal analysis. Li Li: Conceptualization,
480 Methodology, Writing-review & editing.

481

482 *Competing interest.* The authors declare that they have no known competing financial interests
483 or personal relationships that could have appeared to influence the work reported in this paper.

484

485 *Acknowledgements.* This study is supported by the National Natural Science Foundation of
486 China (No.4185161), Shanghai International Science and Technology Cooperation Fund (No.
487 19230742500), and Shanghai Science and Technology Fund (No. 19DZ1205007). Y. Wang
488 was supported by the National Science Foundation. We thank Shanghai Environmental
489 Monitoring Center (SEMC) for conducting the measurement and sharing the data.

490

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492 *Financial support.* This study was financially supported by the National Natural Science
493 Foundation of China (NO. 41875161; NO.42075144), Shanghai International Science and
494 Technology Cooperation Fund (NO. 19230742500), Shanghai Science and Technology Fund
495 (No. 19DZ1205007), Shanghai Sail Program (NO.19YF1415600), and the National Key

496 Research and Development Program of China (NO.2018YFC0213600). Y. Wang was
497 supported by the National Science Foundation.

498

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