



- 1 A measurement and model study on ozone characteristics in marine air at a remote
- 2 island station and its interaction with urban ozone air quality in Shanghai, China
- 3 Yixuan Gu^{a,b}, Fengxia Yan^c, Jianming Xu^{a,b,*}, Yuanhao Qu^{a,b}, Wei Gao^{a,b}
- 4 ^aShanghai Typhoon Institute, Shanghai Meteorological Service, Shanghai 200030, China
- 5 ^bShanghai Key Laboratory of Meteorology and Health, Shanghai Meteorological Service,
- 6 Shanghai 200030, China
- 7 ^cMeteorological Center of Traffic Management of East China, Shanghai 2000135, China
- 8 Corresponding to: Dr. Jianming Xu (metxujm@163.com)
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11 Abstract

12	To understand the characteristics and changes of baseline ozone (O_3) in oceanic air in
13	East China, a six-year measurement of O_3 concentration was conducted from January 1
14	2012 to September 15 2017 at a remote offshore station located on the Sheshan Island
15	(SSI) near the megacity of Shanghai. The observed monthly mean $\rm O_3$ concentrations at
16	SSI ranged from 33.4 to 61.4 ppbv during the study period, which were about 80% and 12%
17	higher, respectively than those measured at downtown and rural sites in Shanghai.
18	Compared to the remarkable O_3 increases observed at urban and rural sites in Shanghai,
19	observed O_3 concentrations at SSI exhibited statistically insignificant increasing changes
20	(1.12 ppbv yr ⁻¹ , α >0.10) during the observation period, suggesting less impacts of
21	anthropogenic emissions on O_3 levels in oceanic air. In addition, an insignificant
22	decreasing change (-0.72 ppbv yr $^{\text{-1}},$ $\alpha\text{>}0.10)$ was detected in O_3 concentrations at SSI in
23	September and October when the influence of regional transport was minimum
24	throughout the year, providing a good proxy to study the baseline oxidation capacity of the
25	oceanic atmosphere. City plumes from Shanghai usually carried higher levels of $NO_{x},$
26	resulting in decreased O_3 concentrations at SSI during southwesterly and westerly winds.
27	However, In MAM (March–May) and JJA (June–August), due to the enhanced oxidation of
28	oxygenated volatile organic compounds, O_3 could be continuously produced during
29	daytime in aged city plumes, resulting in elevated O ₃ concentrations transported to SSI.
30	The impacts of the offshore O_3 on O_3 levels in Shanghai are quantified during an easterly
31	wind dominant episode (September 1-30, 2014) using the WRF-Chem model. Sensitivity
32	results suggest that O_3 in the oceanic air inflows can lead to 20–30% increases in urban





- 33 O₃ concentrations, which should be crucially considered in dealing with urban O₃ pollution
- 34 in large coastal cities like Shanghai.





35 1 Introduction

36	Ground-level ozone (O ₃) is a harmful photochemical oxidant detrimental to air quality,
37	human health and land ecosystems (Yue and Unger 2014; Monks et al., 2015; Li et al.,
38	2019a). High ambient ${\rm O}_3$ has been proved to increase the risks of respiratory and
39	cardiovascular mortality (Goodman et al., 2015) and enhance the greenhouse effect
40	(IPCC, 2013). In recent years, O_3 pollution has drawn increasing attention in China, since
41	\ensuremath{O}_3 pollution is getting worse in spite of the implementation of Chinese Clean Air Action
42	Plan. In contrast to the 28-40% decreases in $PM_{2.5}$ (fine particulate matter; diameter ≤ 2.5
43	$\mu m)$ levels, the observed daily maximum 8-h average (MDA8) O_3 concentrations show
44	increasing rates of 1–3 ppb yr ⁻¹ in summer in megacities over eastern China during 2013–
45	2017 (Li et al., 2019b). To address the underlying causes of the increasing $O_{\rm 3}$ pollution
46	has become an urgent issue that triggers lots of discussions based on observational and
47	model studies worldwide.

Observational and model studies indicated that the elevated O3 levels in urban and 48 49 rural areas in eastern China were strongly related to the changes in anthropogenic emissions of O_3 precursors (Ma et al., 2016; Lu et al., 2018; Li et al., 2019b; Gu et al., 50 51 2020). Since the O₃ formation was reported to be under volatile organic compound (VOC) 52 limited regime in most Chinese megacities (e.g. Beijing, Shanghai, and Guangzhou), the sharp decreases in nitrogen oxides (NOx=NO+NO2) emissions combined with slight 53 increases in VOC levels were suggested to be main causes of the observed enhancement 54 of O₃ concentrations in East China (Gao et al., 2017; Xu et al., 2019). In remote areas, 55 changes of baseline O3 also exhibit sensitive responses to human activities (Vingarzan, 56





57	2004; Meng et al., 2009; Wang et al., 2009; Lin et al., 2015). Based on 14-year
58	observations, Wang et al. (2009) pointed out that enhanced pollution flow from the upwind
59	coastal regions contributed to most of the observed increases in O_3 concentrations during
60	1994–2007. And the increase in background O_3 likely made a strong contribution (81%) to
61	the increasing rate of O_3 in urban Hong Kong. To understand the background O_3 changes
62	and its response to human activities are thus necessary in developing long-term
63	strategies to mitigate local O_3 pollution. However, compared to the intensive field studies
64	in polluted cities and surrounding rural regions, continuous observations of O_3 at
65	representative background sites are relatively limited in China (Wang et al., 2017).

66 To better understand the characteristics of the background O₃ changes in mainland China, the China Meteorological Administration (CMA) started to conduct continuous 67 68 measurements of surface O₃ at several regional background stations (e.g. Shangdianzi, Linan, and Longfengshan) since 2005. Over 10-year records from the three sites and 69 Waliguan, a baseline Global Atmospheric Watch (GAW) station in Tibetan Plateau region, 70 71 exhibited different increases in background continental O3 concentrations especially 72 during daytime in China (Lin et al., 2008; Xu et al., 2008; Meng et al., 2009; Ma et al., 73 2016; Xu et al., 2016). The detected positive trends of surface O₃ were in a range of 0.24-74 1.13 ppbv yr⁻¹, suggesting enhanced atmospheric oxidation capacity response to the rapid 75 development of urbanization and industrialization in the past decades. In addition to the 76 changes in background O₃ in terrestrial stations mentioned above, the characteristics of baseline O_3 at remote marine sites are also important. It is because that large amounts of 77 O₃ pollution events occurred in coastal urban agglomerations in East China (Lu et al., 78





79	2018; Li et al., 2019a, b), affected by both city plumes and oceanic air inflows (Tie et al.,
80	2009; Shan et al., 2016). For example, model work of Tie et al. (2009) suggested that sea
81	air masses carried by oceanic inshore air flows aggravated urban O_3 pollution in Shanghai
82	under convergence conditions. Understanding the O_3 characteristics in offshore oceanic
83	regions is therefore an important prerequisite for understanding the land-sea \ensuremath{O}_3
84	interactions and its impacts on O_3 pollution in coastal cities. However, to our knowledge,
85	studies on the characteristics and changes of O_3 in marine air are quite limited in mainland
86	China since it is very difficult to conduct systematic and continuous observations under
87	remote oceanic air conditions.

88 In this report, we present the first relatively long and continuous measurements of O_3 conducted on a remote offshore island (Sheshan Island, SSI) from Jan 2012 to Sep 2017 89 90 in eastern China. The SSI is located at the confluence of the Yellow Sea and the East China Sea, covering about 0.4 square kilometers area. Since there are no inhabitants in 91 92 the island, the observed O3 is seldom affected by local anthropogenic emissions. The 93 collected O₃ data are used to understand the magnitudes and variabilities of O₃ in the 94 offshore regions and their impacts on the O3 concentrations in coastal city areas. First 95 shown are the general impacts of regional transport on the remote atmosphere over the 96 SSI region. Then the diurnal patterns of O3 at SSI are investigated by comparing them with those observed at a downtown site (XJH) in Shanghai. Multi-year changes of O3 97 concentrations at SSI are analyzed to examine the overall changes of baseline O3 in 98 marine air and possible causes. Also analyzed are the impacts of urban plumes on O_3 99 levels in oceanic air in offshore regions. At last, the influence of O3 carried by oceanic air 100





- 101 inflows on urban O₃ air quality in Shanghai is assessed using the Weather Research and
- 102 Forecasting model coupled with Chemistry (WRF-Chem).
- 103 2 Material and methods
- 104 2.1 The SSI site and ozone observations

105 To investigate the characteristics and variabilities of O₃ in marine air and their interactions with urban air quality in coastal areas, ground O3 concentrations were continuously 106 107 measured at SSI site (31.4°N, 122.3°E, 73.5 m a.s.l.), which is approximately 75 108 kilometers away from the east edge of Shanghai city. Figure 1 shows the location of SSI 109 and the surrounding environment. As mentioned in Sect. 1, there is no resident and tourist on the island. The observed O₃ at SSI site can represent the background O₃ conditions in 110 111 oceanic air which are seldom contaminated by anthropogenic emissions. Hourly O_3 data 112 was collected during January 1 2012 to September 15 2017, with a capture rate of 89.7%. O₃ was measured using an analyzer from Ecotech, Australia (Model EC9810), which 113 combined microprocessor control with ultraviolet photometry. The instrument met the 114 technical specifications for United States Environmental Protection Agency, with a quality 115 116 control check every 3 days, filer replaced every 2 weeks and calibration every month. 2.2 Observational data at urban and rural sites in Shanghai 117

- 118 To better understand the characteristics of the offshore O_3 in oceanic air at SSI, O_3
- observations obtained from a downtown site, Xujiahui (XJH) are used for comparisons.
 The XJH site is located at downtown Shanghai, approximately 80 km west from the SSI. In
 addition, since measurements of NO_x, carbon monoxide (CO) and meteorological
 parameters (e.g. wind direction and wind speed) were unavailable at SSI, observations





123	obtained at an adjacent site, Dongtan (DT), are substituted for the investigation. The DT
124	site was set up in a national nature reserve near the coast of Shanghai, where the
125	observed pollutant levels have been reported to well reflect the impacts of pollution
126	transport during the MIRAGE-Shanghai (Megacities Impact on Regional and Global
127	Environment at Shanghai) field campaign (Tie et al., 2013). Similar to SSI, the DT site is
128	also little affected by human activities. The obtained observations of meteorology and
129	pollutants are therefore applied for analyzing the impacts of regional transport on
130	observed O_3 concentrations at SSI. NO_x concentrations were measured with a
131	chemiluminescent trace level analyzer (TEI; Model 42iTL), with detection limit of 0.025
132	ppb. CO concentrations were measured by the Model 48iTL Enhanced CO analyzer,
133	based on gas filter correlation technology. The wind speed and wind direction were
134	measured by using a DZZ4 Automatic Weather Station certificated by the China
135	Meteorological Administration. The geographical locations and surrounding environment
136	of XJH, DT, and SSI are displayed in Fig. 1.

137 2.3 The WRF-Chem model

We simulate O₃ using the regional chemical transport model WRF-Chem (version 3.8, https://www2.acom.ucar.edu/wrf-chem), collaboratively developed through efforts of several institutes, such as the National Center for Atmospheric Research (NCAR) and the National Oceanic and Atmospheric Administration (the National Centers for Environmental Prediction (NCEP). The model includes on-line calculation of meteorological parameters, transport, mixing, emission, and chemical transformation of trace gases and aerosols (Grell et al., 2005). The Regional Acid Deposition Model version 2 (RADM2, Stockwell et





145	al., 1990) gas-phase chemical mechanism is used for the O_3 formation chemistry.
146	Photolysis rates are calculated by using the fast radiation transfer module (FTUV)
147	followed those in Madronich and Flocke (1999) and Tie et al. (2003). ISORROPIA II
148	secondary inorganic (Fountoukis and Nenes, 2007) and the Secondary ORGanic Aerosol
149	Model (SORGAM) (Schell et al., 2001) schemes are used for aerosol chemistry. Dry
150	deposition follows the standard resistance-in-series model of Wesely (1989). The major
151	physical processes employed in the model follow the Lin microphysics scheme (Lin et al.,
152	1983), the Yonsei University (YSU) planetary boundary layer (PBL) scheme (Hong and
153	Lim, 2006), the Noah Land surface model (Chen and Dudhia, 2001), and the long-wave
154	radiation parameterization (Dudhia, 1989).

The model used in this study has a horizontal resolution of 6km×6km, including 150 155 156 un-staggered grids in west-east, 150 un-staggered grids in south-north, and 35 vertical layers extending from the surface to 50 hPa. The domain encompasses Shanghai and 157 surrounding region, centered at 31.3°N, 121.4°E. The NCEP FNL (Final) Operational 158 159 Global Analysis data are used for meteorological initial and boundary conditions, with 160 lateral meteorological boundary updated every 6 h. Basic chemical lateral boundary 161 conditions are constrained by a global chemical transport model (MOZART-4, Model for 162 OZone And Related chemical Tracers, version 4) (Tie et al., 2001; Emmons et al., 2010). 163 Anthropogenic emissions are derived from the Multi-resolution Emission Inventory for China (MEIC inventory, http://www.meicmodel.org/; Li et al., 2014) for year 2010. Biogenic 164 emissions are calculated online using model of emissions of gases and aerosols from 165 nature (MEGAN2, Guenther et al., 2006). 166





167	2.4 Methods for assessing the trend of ozone
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168	The daily mean O_3 concentrations are used to examine the overall changes in O_3
169	concentrations during the period 2012-2017, including all time of day with qualified
170	measurements. The trends are assessed using two nonparametric methods, which are
171	commonly used to detect trends of non-normally distributed data with seasonality (Xu et
172	al., 2016).The Mann-Kendall (MK) trend test (Mann, 1945; Kendall, 1975; Gilbert, 1987) is
173	used to examine the trend significance, and the Theil-Sen trend estimate method (Sen,
174	1968) is used to estimate the slope of trend, which could also be considered as the rate of
175	change, during the six-year period. Compared to the linear fitting analysis which require
176	data to be independent and follow a Gaussian distribution, the non-parametric trend test
177	methods only need the data be independent (Gocic and Trajkovic, 2013). To determine if
178	the calculated rate of change is statistically significant, the confidence level of at least 95%
179	is adopted in the MK trend test, with α value less than 0.05 being considered a statistically
180	significant trend. The trend significance is examined by comparing the value of a
181	standardized test statistic Z to that of a standard normal variate at a given significance
182	level (Z _a , α =0.05). If Z > Z _{1-$\alpha/2$} , then the dataset is non-stationary, exhibiting either an
183	increasing or a declining trend; If $ Z \leq Z_{1-\alpha/2}$, then the dataset is stationary with no
184	significant trend. Detailed calculation of Z can be referred to Xu et al. (2016).

185 3 Results and discussion

186 3.1 Regional transport characteristics at SSI

187 The observed O_3 concentrations at SSI were inevitably influenced by regional transport

188 depending on the prevailing winds in various seasons. Figure 2 displays the monthly wind





189	rose diagrams averaged over the period of 2012 to 2017 at DT. As mentioned in Sect. 2.2,
190	the DT site is a rural site located quite close to SSI. The observed wind speeds and wind
191	directions could then be applied to deduce the origins of the air mass arriving at SSI in
192	adjacent region. Generally, observed prevailing winds exhibited distinct seasonal
193	variabilities which were greatly affected by the East Asian monsoon. In warm seasons
194	(May-August), the site was predominately influenced by easterly and southeasterly winds,
195	accounting for 40-50% of the total winds. While in cold seasons (November-February),
196	the northwesterly and northerly winds became the predominant flows that affected SSI,
197	accounting for about 45% of total winds. During transitional months (e.g. March, April,
198	September and October), the dominant winds presented more diversities, with wind
199	directions dispersedly distributed in all the directions. The observed seasonal variations of
200	prevailing winds are typical at coastal cities at mid-latitude region (Shan et al., 2016; Xu et
201	al., 2019), suggesting that air masses arriving at SSI originated from various regions and
202	could result in different impacts on the offshore atmospheric composition in different
203	months.

Since CO has a relative long chemical lifetime of a few months, the observed CO concentrations at DT could be regarded as a consequence of regional transport from polluted regions (Tie et al., 2009). Figure 3 displays the observed monthly mean CO mixing ratios under wind directions of north (N), northeast (NE), east (E), southeast (SE), south (S), southwest (SW), west (W), and northwest (NW) at DT during the 2012–2017 period. Observed CO exhibited relative higher concentrations under SW and W winds in all months, with mean mixing ratios of 0.44 and 0.56 ppmv, respectively during 2012–2017





211	(Table 1). The observed high CO mixing ratios suggested that the atmosphere
212	constituents at SSI could be more affected by regional transport of air pollutants under
213	SW and W wind conditions. As SSI is located to the northeast of the Shanghai city (Fig. 1),
214	air masses carried by the SW and W flows usually contain more urban pollutants from
215	upwind city areas, and those carried by E, SE, and NE flows mostly come from the ocean.
216	The oceanic air masses are cleaner compared to those from the cities, leading to lower
217	CO mixing ratio at SSI. For example, observed CO exhibited a mean concentration of
218	0.23 ppmv under SE wind conditions, which was about 50% lower than that influenced by
219	W winds. To further examine the impacts of the SW and W winds on the atmosphere
220	constituents at SSI, Table 2 lists the calculated monthly mean occurrence frequency of the
221	SW and W winds in separate months during the studied period. The SW and S winds
222	were most infrequent in September (6.1 %) and October (5.2 %), suggesting that the
223	atmosphere at SSI during the two months could be less contaminated by pollutants
224	transported from the city and might be more close to the baseline oceanic air conditions.

225 3.2 The diurnal pattern of ozone at SSI

Figure 4 displays the monthly mean diurnal variations of O_3 at SSI and XJH in different months during 2012–2017. The observed O_3 concentrations at the two sites exhibited similar seasonal distinctions, with monthly mean values highest (61.4 ppbv for SSI and 35.9 ppbv for XJH) in May and lowest (33.4 ppbv for SSI and 12.5 ppbv for XJH) in December. Since the O_3 formation in urban Shanghai is VOC-limited, observed O_3 could be significantly depressed by large NO_x emissions at downtown site (XJH) (Gu et al., 2020). In Fig. 4, observed O_3 levels at XJH were quite lower than those at SSI in all





233	months, with mean concentrations of 27.8 and 50.1 ppbv, respectively at XJH and SSI
234	during the observation period. The observed mean O_3 concentration at the rural site of DT
235	(44.7 ppbv) was also lower than that at SSI, which suggested that the O_3 levels in marine
236	air were higher than the continental O_3 at both urban and rural sites. The observed diurnal
237	patterns of O_3 at SSI and XJH were similar to those reported for other sites in eastern
238	China (Xu et al., 2008; Geng et al., 2015; Gao et al., 2017), exhibiting minimums in early
239	morning (06:00–08:00 LST) and maximums in the afternoon (13:00–15:00 LST). However
240	compared to those at the urban site (XJH), observed amplitudes of O_3 diurnal variations
241	were much smaller at SSI. The diurnal variations of surface O_3 can be mainly attributed to
242	the O_3 production through photochemical reactions in the daytime and O_3 depression via
243	NO titration at nighttime (Sillman, 2003). Due to few emissions of O_3 precursors (NO _x and
244	VOCs), the O_3 production and depression could be weaker at remote site, resulting in
245	flatter diurnal cycle of O_3 compared to that at polluted urban site.

Since the amplitudes of O3 diurnal variations usually exhibited much smaller values in 246 247 background areas compared to those in polluted urban regions, the ratio of daily maximum O₃ concentration (O_{3-max}) to minimum O₃ concentration (O_{3-min}) was regarded as 248 249 an indicator to identify if the local O3 pollution was significantly influenced by 250 anthropogenic emissions (Cvitas and Klasinc 1993; Vingarzan, 2004). The O3-max/O3-min 251 ratio displayed larger values in polluted regions (Cvitas et al., 1995) and lower values in less contaminated rural regions. A ratio of about 1.4 suggested that the site could be 252 regarded as a typical background site (Scheel et al., 1997). In Fig. 4, observed O₃ 253 displayed different diurnal variabilities in different months, indicating different influence of 254





255	regional transport on O_3 levels at SSI. Figure 5 displays the calculated monthly mean
256	$O_{3\text{-max}}/O_{3\text{-min}}$ ratios at SSI and XJH, respectively during 2012–2017. Generally, the
257	observed ratios of $O_{3\text{-max}}/O_{3\text{-min}}$ at SSI were much lower than those at XJH in all the
258	months, suggesting less impact of anthropogenic emissions on O_3 levels. The calculated
259	mean ratios were 3.03 and 5.20, respectively at SSI and XJH, and most of the calculated
260	values were larger than 4.50 at the urban site. Besides, the ratios presented distinct
261	seasonal differences at XJH and SSI sites. Higher values were observed in summer,
262	indicating stronger photochemical production of daytime O_3 during June to August. At SSI,
263	the $O_{3\text{-max}}/O_{3\text{-min}}$ ratio exhibited relatively low values in September and October, ranging
264	from 1.61-2.35 during the studied period. Since the observed temperature and solar
265	radiation still exhibited higher values during the two months in Shanghai (Gao et al., 2017),
266	the observed low O_3 diurnal amplitudes should not be attributed to the weakened
267	photochemical formation of O_3 as those in winter. Due to the persistent control of
268	anticyclone, Shanghai and its neighboring areas are usually dominated by stable weather
269	conditions in September and October, resulting in more gentle and diversified wind
270	conditions. During the two months, the occurrences of more polluted SW and W winds
271	were lowest (6.1% and 5.2%) throughout the year. The corresponding wind speed (2.49
272	and 2.50 m $\rm s^{\text{-1}})$ also exhibited values 20% lower those in other months (Table 2). The
273	transport conditions led to fewer pollutants transported to the SSI region, which could
274	explain the observed weak diurnal variabilities of O_3 in September and October. The
275	transport conditions together with O_3 response further confirmed that the transport of city
276	pollutants had minimum impacts on the offshore O_3 levels in oceanic air at SSI in





- 277 September and October.
- 278 3.3 Overall changes of ozone in oceanic air at SSI

279	Several studies have observed increasing trends of ground-level O_3 in metropolitan areas
280	over eastern China since 2013, suggesting that the O_3 increases were mostly attributed to
281	the NO_x emission reductions (Ma et al., 2016; Gao et al., 2017; Lu et al., 2018; Li et al.,
282	2019b). However, the O_3 changes at remote sites were relatively not well elucidated
283	during past years. Figure 6a presents the monthly variations of O_3 concentrations at SSI
284	and XJH during the 2012-2017 period. The statistical results of the MK test and Theil-
285	Sen trend estimate method indicated that observed monthly mean O_3 mixing ratios (O_{3-ave})
286	exhibited increasing changes at both urban (XJH) and remote sites (SSI) in Shanghai,
287	with calculated increasing rate of 1.97 and 1.12 ppbv yr ⁻¹ , respectively in XJH and SSI.
288	Though an overall upward trend of O_3 was detected at SSI, the changes were not as
289	remarkable as those observed at XJH, which could not even pass the MK trend test at the
290	90% confidence level. The detected increasing trend of O_3 in oceanic air at SSI were also
291	not as remarkable as those observed at remote continental sites (e.g. DT and Lin'an) over
292	the Yangtze River Delta (YRD) region (Xu et al., 2008; Gao et al., 2017; Gu et al., 2020),
293	suggesting few influence of anthropogenic emissions on the observed O_3 levels in
294	oceanic air.
205	As discussed in Sect 3.1 different proveiling winds led to different transport

As discussed in Sect. 3.1, different prevailing winds led to different transport conditions at SSI in various months. Comparatively, the observed O_3 concentrations at SSI were least contaminated by the regional transport of air pollutants in September and October. To further examine the changes of nearly uncontaminated O_3 , we present the





299	variations of daily mean surface O_3 mixing ratios in September and October at SSI and
300	XJH, respectively during the six-year period in Fig. 6b. The corresponding mean O_3
301	mixing ratios were 60.9 and 31.3 ppbv, respectively at SSI and XJH. Compared to the
302	significant increasing changes of O ₃ (0.59 ppbv yr ⁻¹ , α <0.05) at urban site (XJH), observed
303	O_3 at SSI in September and October exhibited insignificant decreasing changes during
304	the studied period, with an average rate of -0.72 ppbv yr ⁻¹ . The changes were smaller and
305	quite different from the overall changes at SSI as well as those detected at XJH, which
306	further indicated that the observed O_3 levels at SSI in September and October should be
307	seldom influenced by urban plumes, providing a good proxy to study the baseline O_3 and
308	oxidation capacity of background atmosphere in eastern China.

309 To investigate possible drivers of the observed changes in the least contaminated O₃ 310 concentrations in September and October at SSI, Table 3 displays the statistical results of the MK test and Theil-Sen trend estimate for NO_x, CO mixing ratios, temperature, and 311 wind speed during the 2012-2017 period. Statistically significant upward trends were 312 313 detected in NOx concentrations and wind speed with estimated increasing rates of 0.48 ppbv yr⁻¹ and 0.21 m s⁻¹ yr⁻¹, respectively during the observation period (α <0.05). The 314 315 results suggested that increases in surface wind speed might be an important 316 meteorological driver of the observed decreasing changes in O3 levels at SSI from 2012 to 317 2017. Since both NOx and CO levels exhibited different increases, it means more pollutants could be transported to the island and might resulted in elevated O3 318 depressions during the period. Figure 6c presents corresponding variations of daytime 319 320 and nighttime mean O₃ concentrations at SSI. Similarly, insignificant downward changes

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322	depression of O_3 might be enhanced at SSI due to elevated wind speeds and NO_x
323	concentrations. Since observations of solar radiation were not available during the study
324	period, the influence of radiation cannot be analyzed which might also play important role
325	in determining the overall changes of the observed O_3 at SSI.
326	3.4 Impacts of urban plumes on ozone in oceanic air at SSI
327	Due to the relatively long residence lifetime (about one month), O_3 produced at urban
328	regions could be transported several hundred kilometers away to downwind areas.
329	Meanwhile, the urban plumes become more aged with continuous production/depletion of
330	O_3 and its precursors, resulting in non-linear variations of O_3 in downwind areas (Geng et

were detected in both daytime and nighttime O₃ levels, indicating that the diffusion and

al., 2011; Tie et al., 2009, 2013). Several studies suggested that there tended to be 331 332 considerable O₃ formations in aged urban plumes in the downwind region of Shanghai (Geng et al., 2011; Tie et al., 2013). To investigate the impacts of urban plumes on the O_3 333 levels in oceanic air at SSI, the relationships between observed O_3 and NO_x under 334 335 different wind conditions at SSI and DT are investigated in this section.

336 Figure 7 presents the daytime and nighttime O3/NOx-wind relationships in MAM 337 (March-May), JJA (June-August), SON (September-November), and DJF (December-338 February), respectively during 2012-2017. The SW and W winds were associated with 339 higher NO_x concentrations in both daytime and nighttime. The result is consistent with the observed CO changes in Sect. 3.1. Since there is no local anthropogenic emission at SSI, 340 the higher levels of NO_x and CO were mainly resulted from the transport of more polluted 341 urban plumes by the SW and W winds. Generally, observed daytime O_3 and NO_x 342





343	concentrations presented opposite variations with the wind direction changes (Fig. 7a). In
344	SON and DJF, the correlation coefficients (Rs) between daytime O_3 and NO_x were -0.72
345	and -0.75, respectively, indicating that the O_3 formation was inhibited by increased NO_x
346	concentrations. The results are in accordance with Tie et al. (2013), who suggested that
347	the VOC-limited regime of O_3 formation was not only confined in urban Shanghai, but also
348	extended to a broader regional area surrounding Shanghai. However, in MAM and JJA,
349	the daytime $O_3\text{-}NO_x$ variations presented totally different patterns under SW and W wind
350	conditions. As wind directions turned from E-SE to SW-W, observed mean $\ensuremath{NO_{x}}$
351	concentrations increased from about 10 ppbv to 20 ppbv, while observed mean \ensuremath{O}_3
352	concentrations increased from 50-60 ppbv to 70-80 ppbv. The enhancements in daytime
353	O_3 levels suggested that there should be persistent production of O_3 in the polluted air
354	masses carried by the SW and W winds in MAM and JJA.

Based on observations and WRF-Chem simulations, Tie et al. (2013) suggested 355 considerable O3 production in aged city plumes in the downwind area of Shanghai. Since 356 357 air masses affecting SSI site were directly originated from Shanghai under the SW and W wind conditions (Fig. 1), the observed O₃ enhancements should be mainly attributed to the 358 359 O_3 production in the city plumes carried by SW and W winds. Studies during the 360 MIRAGE-Shanghai campaign suggested several factors that contributed to the O3 361 enhancements in aged city plumes downwind Shanghai. First, as there is a large area of forest located in the south of Shanghai, Geng et al. (2011) suggested that continuous 362 oxidation of isoprene emitted by the biogenic sources could result in enhanced production 363 364 of hydrogen radicals (HO₂) especially in warm seasons. Once the air massed were





365	transported north and mixed with high NO_x emissions, O_3 would be quickly produced.
366	However, the impacts of biogenic emissions on O_3 production were mainly limited in the
367	south part of Shanghai, which can hardly influence the atmosphere in the SSI region.
368	Then, Tie et al. (2013) further illustrated that the OH reactivity of alkane, alkene, aromatics,
369	and oxygenated VOCs (OVOCs) contributed to the \ensuremath{O}_3 formation in city plumes. Among
370	them, the influence of alkane, alkene and aromatics mostly occurred within or near the city,
371	while the OVOCs could be produced or emitted during the transport of the city plumes,
372	resulting in substantial O_3 enhancements in aged city plumes at 100–200 km downwind
373	Shanghai.
374	The SSI is located approximately 100 km northeast from the downtown area of
375	Shanghai. In MAM and JJA, the SW and W winds carried air masses with enhanced
376	OVOCs oxidation and O_3 production, resulting in elevated daytime O_3 levels on the island.
377	While in SON and DJF, the observed O_3 decreases at SSI during SW and W winds
378	suggested lower efficiency of O_3 productivity in the city plumes. That might be because
379	that fewer OVOCs were released or produced downwind the city due to the lower
380	temperature and weaker solar radiation (Cai et al., 2009). In addition, in SON and DJF, the

SW and W winds were usually related to low pressure system with large cloud cover and rich water vapor in Shanghai, which could also lead to depressed photochemical reactions and decreased O₃ levels. At night, observed O₃ and NO_x displayed totally opposite changes with wind directions (Fig. 7b), indicating O₃ depression by nighttime NO_x titration in all the seasons. High O₃ levels were observed under northeasterly, easterly and southeasterly oceanic wind conditions, ranging from 50–60, 30-55, 55–60, and 40–50





- 387 ppbv respectively at night in MAM, JJA, SON, and DJF. More detailed measurements are
- 388 still needed to further understand the impacts of city plumes on the O₃ levels in oceanic
- 389 air.
- 390 3.5 Impacts of offshore ozone on urban ozone air quality in Shanghai
- 391 As is presented in Sect. 3.2 and 3.3, observed O₃ concentrations at SSI were much higher
- than those at urban site (XJH), suggesting higher levels of O₃ in oceanic air than those on
- 393 the continent. Therefore, sea breezes tend to bring more O_3 to the continent, aggravating
- 394 O₃ pollution in coastal cities. Shanghai is one of the largest cities located on the east coast
- of China, experiencing severe O_3 pollution in recent years (Xu et al., 2019; Gu et al.,
- 396 2020). The easterly winds from the ocean greatly affect the Shanghai region, accounting
- 397 for 64-78 % of the total flows in non-winter months during the period 2012-2017. To
- 398 understand the impacts of higher O_3 in oceanic air on the urban air quality, numerical
- 399 experiments are conducted using the WRF-Chem model to examine the response of O_3
- 400 levels in Shanghai to various oceanic air inflow conditions in this section.

401 Simulations are performed during September 1-30 2014 when the prevailing winds 402 were mostly northeasterly and easterly in the Shanghai region. The occurrence 403 frequencies of the northeasterly and easterly winds were 23% and 27% respectively, 404 during the simulation period, suggesting dominant influence of the oceanic air inflows on the city of Shanghai. Consistent with above analysis, observed air O₃ concentrations were 405 406 much higher in oceanic regions than those in city areas, with monthly mean values of 30.9 407 and 57.7 ppbv, respectively at XJH and SSI in September 2014. The chemical boundary 408 conditions (BCs) of the regional model can represent the inflows conditions to explore





409	their impacts on surface concentrations of air pollutants over a certain continent region.
410	Using this method, Pfiter et al. (2011) proposed that chemical inflows taken from different
411	observational and model datasets could result in differences of ± 15 ppbv in O ₃ levels in
412	the US west coast region. Therefore, three sets of numerical experiments are conducted
413	as follows to access the impacts of oceanic O_3 air inflows on the urban O_3 air quality in
414	Shanghai. All the simulations are driven by the same emissions, initial conditions, physical
415	and chemical schemes.
416	(1) BC_40: O_3 concentrations at the eastern lateral boundary of the domain on the
417	ocean are assigned to 40 ppbv, which is provided by the MOZART-4 model, closed to the
418	observed urban O_3 levels (29.0–38.4 ppbv) in Shanghai in September. The chemical BCs
419	are updated every 6 hours.
420	(2) BC_50: Same as BC_40, but with O_3 concentrations setting to 50 ppbv at the
421	eastern lateral boundary of the domain.
422	(3) BC_60: Same as BC_40, but with O_3 concentrations at the eastern lateral
423	boundary of the domain setting to 60 ppbv according to the observed O_3 levels at SSI
424	(50.9–71.0 ppbv) in September.
425	Figure 8 displays the simulated monthly mean distributions of surface O_3
426	concentrations in BC_40, BC_50 and BC_60 scenarios, respectively. The calculated
427	distributions of O_3 agree with observations, which exhibit lower values in urban regions
428	compared to those in rural and ocean areas, indicating strong O_3 depressions in the city of
429	Shanghai due to the VOC-limited ${\rm O}_3$ formation regime. The R values between the
430	simulated and observed O_3 concentrations are all larger than 0.50 at XJH, suggesting





431	good prediction of O ₃ variations by the model. Table 4 displays the statistical results of the
432	comparisons between the simulated and observed surface O_3 concentrations at SSI and
433	XJH, respectively. Generally, the WRF-Chem model underestimates O_3 concentrations at
434	both XJH and SSI site. Taken the BC_40 scenario for example, the O_3 concentrations are
435	underestimated by 27% and 36% respectively at XJH and SSI, suggesting larger
436	underestimation of O_3 concentrations in ocean regions. Model results suggest that
437	increases in the eastern boundary O_3 lead to increases in O_3 concentrations at both urban
438	and remote sites when the prevailing winds are mostly easterly in Shanghai. With \ensuremath{O}_3
439	concentrations increasing from 40 to 60 ppbv in the easterly inflows, the simulated
440	monthly mean O_3 concentrations increase by 7.0 and 10.4 ppbv, respectively at XJH and
441	SSI. The underestimation of O_3 levels by the model is also significantly improved in the
442	BC_60 scenario, with the chemical BCs of O_3 more close to the observed chemical inflow
443	conditions. Compared to those in the BC_40 scenario, the normalized mean bias (NMBs)
444	of the predicted O_3 concentrations reduced from -36.1 $\%$ to -18.1 $\%$ at SSI and -27.6% to
445	-4.6% at XJH in the BC_60 scenario, suggesting a crucial role of the eastern oceanic air
446	inflows in influencing O_3 air quality in Shanghai.

The calculated monthly mean differences in surface O_3 concentrations between simulations in different scenarios are further presented in Fig. 9. Since the dominant winds are easterly during the simulation period, distinct changes in surface O_3 concentrations throughout Shanghai are generated, exhibiting generally gradient increases from the ocean to the continent as O_3 increases in the oceanic air inflows. With every 10 ppbv increases in O_3 levels in oceanic air, the simulated surface mean O_3 concentrations





453	increase by 3-6 ppbv in the land area and 4-7 ppbv in the offshore region. Due to the
454	strong O_3 depressions associated with high anthropogenic emissions, the simulated O_3
455	enhancements are relatively lower in the central urban region compared to those in
456	surrounding areas. Even so, simulated mean O_3 concentrations still exhibit 6–8 ppbv
457	increases in the BC_60 scenario, accounting for approximately 30% of the simulated \ensuremath{O}_3
458	concentrations in the BC_40 case. During the period 2012-2017, most of the measured
459	O_3 concentrations ranged between 50—60 ppbv at SSI in non-winter seasons. Carried by
460	the easterly inflows, these oceanic air masses with higher O_3 levels (50—60 ppbv) could
461	be transported to the coastal regions, resulting in approximately 20-30% increases in
462	urban O_3 concentrations in Shanghai according to the sensitivity results.
463	4 Conclusions
464	In this paper, we present the first relatively long and continuous measurements of oceanic
465	
	air O_{3} conducted at an offshore monitoring station on the Sheshan Island during January 1 $% \left(1-\frac{1}{2}\right) =0$
466	air O_3 conducted at an offshore monitoring station on the Sheshan Island during January 1 2012 to September 15 2017. The southwesterly and westerly winds are proved to carry
466 467	air O_3 conducted at an offshore monitoring station on the Sheshan Island during January 1 2012 to September 15 2017. The southwesterly and westerly winds are proved to carry more pollutants to the SSI site, exerting greater influence of human activities on the
466 467 468	air O_3 conducted at an offshore monitoring station on the Sheshan Island during January 1 2012 to September 15 2017. The southwesterly and westerly winds are proved to carry more pollutants to the SSI site, exerting greater influence of human activities on the oceanic atmosphere over the offshore region of the East China Sea. Since the two kinds
466 467 468 469	air O ₃ conducted at an offshore monitoring station on the Sheshan Island during January 1 2012 to September 15 2017. The southwesterly and westerly winds are proved to carry more pollutants to the SSI site, exerting greater influence of human activities on the oceanic atmosphere over the offshore region of the East China Sea. Since the two kinds of winds exhibited minimum occurrence frequencies and wind speeds in September and
466 467 468 469 470	air O ₃ conducted at an offshore monitoring station on the Sheshan Island during January 1 2012 to September 15 2017. The southwesterly and westerly winds are proved to carry more pollutants to the SSI site, exerting greater influence of human activities on the oceanic atmosphere over the offshore region of the East China Sea. Since the two kinds of winds exhibited minimum occurrence frequencies and wind speeds in September and October, atmosphere at SSI during the two months are considered to be less affected by
466 467 468 469 470 471	air O ₃ conducted at an offshore monitoring station on the Sheshan Island during January 1 2012 to September 15 2017. The southwesterly and westerly winds are proved to carry more pollutants to the SSI site, exerting greater influence of human activities on the oceanic atmosphere over the offshore region of the East China Sea. Since the two kinds of winds exhibited minimum occurrence frequencies and wind speeds in September and October, atmosphere at SSI during the two months are considered to be less affected by the transport of regional pollution.
466 467 468 469 470 471 472	air O ₃ conducted at an offshore monitoring station on the Sheshan Island during January 1 2012 to September 15 2017. The southwesterly and westerly winds are proved to carry more pollutants to the SSI site, exerting greater influence of human activities on the oceanic atmosphere over the offshore region of the East China Sea. Since the two kinds of winds exhibited minimum occurrence frequencies and wind speeds in September and October, atmosphere at SSI during the two months are considered to be less affected by the transport of regional pollution. Compared to those in urban (XJH) and rural (DT) sites, the observed O ₃ levels were

474 Similar seasonal and diurnal patterns of O₃ were observed at SSI and XJH; however, the





475	amplitudes of O_3 variations were much smaller at the offshore site (SSI). Since O_3
476	formation in Shanghai and its surrounding regions were VOC-limited, the observational
477	results suggested that the production and depression of O_3 could be weaker in the ocean
478	regions due to weak influence of the anthropogenic emissions. Observed mean
479	O_{3-max}/O_{3-min} ratios also exhibited lower values at SSI (3.03) than those at XJH (5.20), with
480	minimum values ranging from 1.61-2.35 in September and October. The result further
481	illustrated that SSI was seldom affected by the anthropogenic emissions, especially in
482	September and October.

The multi-year changes of the oceanic O3 at SSI are investigated using the 483 484 Mann-Kendall trend test and the Theil-Sen trend estimate method during 2012-2017. Different from the significant O₃ increases detected at XJH and other rural sites reported 485 486 in previous studies, the observed mean O3 concentrations at SSI exhibited statistically insignificant increasing changes (1.12 ppbv yr⁻¹, α>0.10) during the observation period 487 and insignificant decreasing changes (-0.72 ppbv yr⁻¹, α>0.10) in September and October 488 489 when the transport of city pollutants had minimum impacts on the island. Due to fewer 490 impacts of anthropogenic emissions, most of the observed changes in O3 at SSI could be 491 attributed to the changes of meteorological parameters. Observed wind speed exhibited 492 significant increases (0.21 m s⁻¹, α >0.10) in September and October during the observation period, suggesting that enhanced diffusion conditions could be an important 493 meteorological factor in determining the decreases in O_3 concentrations during the 494 495 observation period.

496 The impacts of urban plumes on O_3 levels in oceanic air at SSI are evaluated by





497	studying the relationships between observed O_3 and NO_x under different wind conditions.
498	The SW and W winds usually carried air masses with higher NO_x concentrations in both
499	daytime and nighttime to the island. Generally, observed daytime and nighttime O_3
500	concentration decreased as $\ensuremath{NO_x}$ concentration increases in SW and W winds, exhibiting
501	typical VOC-limited characteristics of O_3 formation. The pattern was more typical in SON
502	and DJF, with R values of -0.72 and -0.75, respectively between O_3 and NO_x
503	concentrations. In MAM and JJA, the daytime $O_3\text{-}NO_x$ variations presented kind of
504	positive relationships under SW and W wind conditions, suggesting continuous O_3
505	production in aged city plumes from Shanghai. As reported in previous studies during the
506	MIRAGE-Shanghai campaign, enhanced OVOCs oxidation should be the most important
507	driver of the observed O_3 enhancements in the city plumes transported by the SW and W
508	winds.
508 509	winds. The influence of the oceanic O_3 air inflows on urban O_3 air quality in Shanghai are
508 509 510	winds. The influence of the oceanic O_3 air inflows on urban O_3 air quality in Shanghai are quantified during an easterly wind dominant episode (September 1–30, 2014). Numerical
508 509 510 511	winds. The influence of the oceanic O ₃ air inflows on urban O ₃ air quality in Shanghai are quantified during an easterly wind dominant episode (September 1–30, 2014). Numerical experiments are conducted with chemical BCs of O ₃ assigned according to different inflow
508 509 510 511 512	winds. The influence of the oceanic O ₃ air inflows on urban O ₃ air quality in Shanghai are quantified during an easterly wind dominant episode (September 1–30, 2014). Numerical experiments are conducted with chemical BCs of O ₃ assigned according to different inflow conditions using the WRF-Chem model. Model results suggest that increases of O ₃ in the
508 509 510 511 512 513	winds. The influence of the oceanic O ₃ air inflows on urban O ₃ air quality in Shanghai are quantified during an easterly wind dominant episode (September 1–30, 2014). Numerical experiments are conducted with chemical BCs of O ₃ assigned according to different inflow conditions using the WRF-Chem model. Model results suggest that increases of O ₃ in the easterly oceanic air inflows will lead to gradient increases from the ocean to the continent.
508 509 510 511 512 513 514	winds. The influence of the oceanic O ₃ air inflows on urban O ₃ air quality in Shanghai are quantified during an easterly wind dominant episode (September 1–30, 2014). Numerical experiments are conducted with chemical BCs of O ₃ assigned according to different inflow conditions using the WRF-Chem model. Model results suggest that increases of O ₃ in the easterly oceanic air inflows will lead to gradient increases from the ocean to the continent. With every 10 ppbv O ₃ increases, the calculated surface mean O ₃ concentrations can
508 509 510 511 512 513 514 515	winds. The influence of the oceanic O ₃ air inflows on urban O ₃ air quality in Shanghai are quantified during an easterly wind dominant episode (September 1–30, 2014). Numerical experiments are conducted with chemical BCs of O ₃ assigned according to different inflow conditions using the WRF-Chem model. Model results suggest that increases of O ₃ in the easterly oceanic air inflows will lead to gradient increases from the ocean to the continent. With every 10 ppbv O ₃ increases, the calculated surface mean O ₃ concentrations can increase by 3–6 ppbv in the land and 4–7 ppbv in the offshore region. Compared to those
508 509 510 511 512 513 514 515 516	winds. The influence of the oceanic O ₃ air inflows on urban O ₃ air quality in Shanghai are quantified during an easterly wind dominant episode (September 1–30, 2014). Numerical experiments are conducted with chemical BCs of O ₃ assigned according to different inflow conditions using the WRF-Chem model. Model results suggest that increases of O ₃ in the easterly oceanic air inflows will lead to gradient increases from the ocean to the continent. With every 10 ppbv O ₃ increases, the calculated surface mean O ₃ concentrations can increase by 3–6 ppbv in the land and 4–7 ppbv in the offshore region. Compared to those in surrounding regions, O ₃ in central city of Shanghai exhibited lower enhancements in
508 509 510 511 512 513 514 515 516 517	winds. The influence of the oceanic O ₃ air inflows on urban O ₃ air quality in Shanghai are quantified during an easterly wind dominant episode (September 1–30, 2014). Numerical experiments are conducted with chemical BCs of O ₃ assigned according to different inflow conditions using the WRF-Chem model. Model results suggest that increases of O ₃ in the easterly oceanic air inflows will lead to gradient increases from the ocean to the continent. With every 10 ppbv O ₃ increases, the calculated surface mean O ₃ concentrations can increase by 3–6 ppbv in the land and 4–7 ppbv in the offshore region. Compared to those in surrounding regions, O ₃ in central city of Shanghai exhibited lower enhancements in response to the O ₃ increases in oceanic air inflows due to strong O ₃ depression processes.





- 519 urban O₃ concentrations which should be crucially considered in dealing with O₃ pollution
- 520 in large coastal cities like Shanghai.
- 521
- 522 Data availability. The data used in this paper can be provided upon request from Dr.
- 523 Jianming Xu (metxujm@163.com).
- 524
- 525 Author contribution. YG and JX came up with the original idea, designed the analysis
- 526 methods, developed the model code, and performed the simulations. WG provided the
- 527 observational data. YG and YQ conducted the analysis of the observations and model
- 528 results. YG prepared the manuscript with contributions from all co-authors.
- 529
- 530 *Competing interest.* The authors declare that they have no conflict of interest.
- 531
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535 Reference

536	Cai, C., Geng, F., Tie, X. X., Yu, Q., and An, J.: Characteristics and source apportionment
537	of VOCs measured in Shanghai, China, Atmos. Environ., 44, 5005–5014, 2010.
538	Chen, F. and Dudhia, J.: Coupling an advanced land surface hydrology model with the
539	Penn State-NCAR MM5 modeling system, Part I: Model implementation and
540	sensitivity, Mon.Weather Rev., 129, 569–585, 2001.
541	Cvitas, T., and Klasinc, L.: Measurement of tropospheric ozone in the Eastern
542	Mediterranean, Boll. Geofisico, 16, 521–527, 1993.
543	Cvitas, T., Kezele, N., Klasinc, L., and Lisac, J.: Tropospheric ozone measurements in
544	Croatia, Pure Appl. Chem., 67, 1450–1453,1995.
545	Dudhia, J.: Numerical study of convection observed during the winter monsoon
546	experiment using a mesoscale two-dimensional model, J. Atmos. Sci., 46, 3077-
547	3107, 1989.
548	Emmons, L. K., Walters, S., Hess, P. G., Lamarque, JF., Pfister, G. G., Fillmore, D.,
549	Granier, C., Guenther, A., Kinnison, D., Laepple, T., Orlando, J., Tie, X., Tyndall,
550	G., Wiedinmyer, C., Baughcum, S. L., and Kloster, S.: Description and evaluation
551	of the Model for Ozone and Related chemical Tracers, version 4 (MOZART-4),
552	Geosci. Model Dev., 3, 43–67, https://doi.org/10.5194/gmd-3-43-2010, 2010.
553	Fountoukis, C., and Nenes, A.: ISORROPIA II: a computationally efficient aerosol
554	thermodynamic equilibrium model for K ⁺ , Ca ²⁺ , Mg ²⁺ , NH ₄ ⁺ , Na ⁺ , SO ₄ ²⁻ , NO ₃ ⁻ ,
555	Cl ⁻ ,H ₂ O aerosols. Atmos. Chem. Phys. 7, 4639–4659, 2007.
556	Gao, W., Tie, X., Xu, J., Huang, R., Mao, X., Zhou, G., and Chang, L.: Long-term trend of
557	O_3 , in a mega city (Shanghai), China: characteristics, causes, and interactions
558	with precursors, Sci. Total Environ., 603–604, 425–433, 2017.
559	Geng, F., Mao, X., Zhou, M., Zhong, S., and Lenschow, D.: Multi-year ozone concentration
560	and its spectra in Shanghai, China, Sci. Total Environ., 521–522, 135–143, 2015.
561	Geng, F., Tie, X., Guenther, A., Li, G., Cao, J., and Harley, P.: Effect of isoprene emissions
562	from major forests on ozone formation in the city of Shanghai, China, Atmos.
563	Chem. Phys., 11, 10449–10459, 2011.
564	Gilbert, R.O.: Statistical Methods for Environmental Pollution Monitoring. John Wiley &
565	Sons, New York, USA, 1987.
566	Goodman, J. E., Prueitt, R. L., Sax, S. N., Pizzurro, D. M., Lynch, H. N., Zu, K., and
567	Venditti, F. J.: Ozone exposure and systemic biomarkers: evaluation of evidence
568	for adverse cardiovascular health impacts, Crit. Rev. Toxicol., 45, 412–452, 2015.
569	Grell, G. A., Peckham, S. E., Schmitz, R., McKeen, S. A., Frost, G., Skamarock, W. C.,
570	and Eder, B.: Fully coupled "online" chemistry within the WRF model, Atmos.
571	Environ., 39, 6957–6975, 2005.
572	Gocic, M., Trajkovic, S.: Analysis of changes in meteorological variables using
573	Mann-Kendall and Sen's slope estimator statistical tests in Serbia, Global Planet.
574	Change, 100, 172–182, 2013.
575	Gu, Y., Li, K., Xu, J., Liao, H., Zhou, G.: Observed dependence of surface ozone on
576	increasing temperature in Shanghai, China. Atmos. Environ., 221, 117108, 2020.
577	Guenther, A., Karl, T., Harley, P., Wiedinmyer, C., Palmer, P. I., and Geron, C.: Estimates





578	of global terrestrial isoprene emissions using MEGAN (model of emissions of
579	gases and aerosols from nature), Atmos. Chem. Phys., 6, 3181–3210, 2006.
580	Hong, S. Y. and Lim, J. O. J.: The WRF Single-Moment 6-Class Microphysics Scheme
581	(WSM6), J. Korean Meteor. Soc., 42, 129–151, 2006.
582	IPCC: Climate Change 2013: The Physical Science Basis. Contribution of Working Group
583	I to the Fifth Assessment Report of the Intergovernmental Panel on Climate
584	Change, edited by Stocker, T. F., Qin, D., Plattner, G. K., Tignor, M., Allen, S. K.,
585	Boschung, J., Nauels, A., Xia, Y., Bex, V. and Midgley, P. M.,, Cambridge
586	University Press, United Kingdom and New York, USA, 2013.
587	Kendall, M.G.: Rank Correlation Methods, fourth ed., Charles Griffin, London, 1975.
588	Li, K., Jacob, D.J., Liao, H., Zhu, J., Shah, V., Shen, L., Bates, K.H., Zhang, Q., and Zhai,
589	S.: A two-pollutant strategy for improving ozone and particulate air quality in China.
590	Nat. Geosci. https://doi.org/10.1038/s41561-019-0464-x, 2019a.
591	Li, K., Jacob, D. J., Liao, H., Shen, L., Zhang, Q., and Bates, K.H.: Anthropogenic drivers
592	of 2013–2017 trends in summer surface ozone in China, P. Natl. A. Sci. USA 116
593	(2), 422–427, 2019b.
594	Li, M., Zhang, Q., Streets, D., He, K. B., Cheng, Y. F., Emmons, L. K., Huo, H., Kang, S. C.,
595	Lu, Z., Shao, M., Su, H., Yu, X., and Zhang, Y.:. Mapping Asian anthropogenic
596	emissions of non-methane volatile organic compounds to multiple chemical
597	mechanisms. Atmos. Chem. Phys. 14, 5617–5638, 2014.
598	Lin, M., Horowitz, L. W., Cooper, O. R., Tarasick, D., Conley, S., Iraci, L. T., Johnson, B.,
599	Leblanc, T., Petropavlovskikh, I., and Yates, E. L: Revisiting the evidence of
600	increasing springtime ozone mixing ratios in the free troposphere over western
601	North America, Geophys. Res. Lett., 42, 8719–8728,
602	https://doi.org/10.1002/2015GL065311, 2015.
603	Lin, Y. L., Farley, R. D., and Orville, H. D.: Bulk parameterization of the snowfield in a
604	cloud model, J. Clim. Appl. Meteorol., 22, 1065–1092, 1983.
605	Lin, W., Xu, X., Zhang, X., Tang, J.:. Contributions of pollutants from North China Plain to
606	surface ozone at the Shangdianzi GAW Station, Atmos. Chem. Phys., 8,5889-
607	5898, 2008.
608	Lu, X., Hong, J., Zhang, L., Cooper, O.R., Schults, M. G., Xu, X., Wang, T., Gao, M., Zhao,
609	Y., and Zhang, Y.: Severe surface ozone pollution in China: a global perspective,
610	Environ. Sci. Technol. Lett.,5(8), 487, 194, 2018.
611	Ma, Z., Xu, J., Quan, W., Zhang, Z., Lin, W., and Xu, X.: Significant increase of surface
612	ozone at a rural site, north of eastern China, Atmos. Chem. Phys. ,16, 3969–3977,
613	2016.
614	Madronich, S., and Flocke, S.: The role of solar radiation in atmospheric chemistry, in:
615	Handbook of Environmental Chemistry, edited by Boule, P. Springer, Heidelberg,
616	1–26, https://doi.org/10.1007/978-3-540-69044-3_1, 1999.
617	Mann, H.B.: Non-parametric tests against trend, Econometrica 13, 163–171, 1945.
618	Meng, Z. Y., Xu, X. B., Yan, P., Ding, G. A., Tang, J., Lin, W. L., Xu, X. D., and Wang, S. F.:
619	Characteristics of trace gaseous pollutants at a regional background station in
620	Northern China, Atmos. Chem. Phys., 9, 927–936,
621	https://doi.org/10.5194/acp-9-927-2009, 2009.





622 623	Monks, P. S., Archibald, A. T., Colette, A., Cooper, O., Coyle, M., Derwent, R., Fowler, D., Granier, C., Law, K.S., Mills, G.E., Stevenson, D.S., Tarasova, O., Thouret, V., von
624	Schneidemesser, E., Sommariva, R., Wild, O., and Williams, M.L.; Tropospheric
625	ozone and its precursors from the urban to the global scale from air guality to
626	short-lived climate forcer. Atmos. Chem. Phys., 15, 8889-8973, 2015.
627	Pfister, G. G., Parrish, D. D., Worden, H., Emmons, L. K., Edwards, D. P., Wiedinmver, C.,
628	Diskin, G. S., Huey, G., Oltmans, S. J., Thouret, V., Weinheimer, A., and Wisthaler,
629	A.: Characterizing summertime chemical boundary conditions for airmasses
630	entering the US West Coast, Atmos, Chem. Phys., 11, 1769–1790, 2011.
631	Scheel, H. E., Aresbough, H., Geiss, H., Gormiscek, B., Granby, K., Haszpra, L., Klasinc,
632	L., Kley, D., Laurila, T., Lindskog, A., Roemer, M., Schmitt, R., Simmond, P.,
633	Solberg, S., and Toupande, G.: On the spatial distribution and seasonal variation
634	of lower tropospheric ozone over Europe, J. Atmos. Chem., 28, 11–28, 1997.
635	Schell, B., Ackermann, I., Hass, H., Binkowski, F. S., and Ebel, A.: Modeling the formation
636	of secondary organic aerosol within a comprehensive air quality model system, J.
637	Geophys. Res., 106, 28275–28293, https://doi.org/10.1029/2001JD000384, 2001.
638	Sen, P. K.: Estimates of the regression coefficient based on Kendall's tau, J. Am. Stat.
639	Assoc., 63, 1379–1389, 1968.
640	Shan, W., Yang, P., Lu, H., Ma, K., and Huang, Z.: Influence of Coastal Wind on Surface
641	Ozone and Nitrogen Oxides in Suburban Shanghai, Asia-Pac. J. Atmos. Sci.,
642	52(5), 451–458, 2016.
643	Sillman, S.: Photochemical Smog: Ozone and its Precursors, in: Handbook of Weather,
644	Climate, and Water, edited by Potter, T. and Bradley, R. R., John Wiley & Sons,
645	New York, USA, 227–242, 2003.
646	Stockwell, W. R., Middleton, P., Chang, J. S., and Tang, X.: The second generation
647	regional acid deposition model chemical mechanism for regional air quality
648	modeling, J. Geophys. ResAtmos., 95, 16343–16367, 1990.
649	Tie, X., Brasseur, G., Emmons, L., Horowitz, I., and Kinnison, D.: Effects of aerosols on
650	tropospheric oxidants: a global model study, J. Geophys. ResAtmos., 106,
651	22931–22964, 2001.
652	Tie, X., Geng, F., Guenther, A., Cao, J., Greenberg, J., Zhang, R., Apel, E., Li, G.,
653	Weinheimer, A., Chen, J., and Cai, C.: Megacity impacts on regional ozone
654	formation: observations and WRF-Chem modeling for the MIRAGE-Shanghai field
655	campaign, Atmos. Chem. Phys., 13, 5655–5669, 2013.
656	Tie, X., Geng, F., Peng, L., Gao, W., and Zhao, C.: Measurement and modeling of O_3
657	variability in Shanghai, China: Application of the WRF-Chem model, Atmos.
658	Environ., 43, 4289–4302, 2009.
659	Tie, X., Madronich, S., Li, G. H., Ying, Z.M., Weinheimer, A., Apel, E., and Campos, T.:
660	Simulation of Mexico City plumes during the MIRAGE-Mex field campaign using
661	the WRF-Chem model, Atmos. Chem. Phys. 9, 4621–4638, 2009.
662	Tie, X., Madronich, S., Walters, S., Rasch, P., and Collins, W.: Effect of clouds on
663	photolysis and oxidants in the troposphere, J. Geophys. Res., 108, 4642, 2003.
664	Vingarzan, R.: A review of surface ozone background levels and trends, Atmos. Environ.,
665	38, 3431–3442, 2004.





666	Wang, T., Wei, X. L., Ding, A. J., Poon, C. N., Lam, Y. S., Li, Y. S., Chan, L. Y., and Anson,
667	M.: Increasing surface ozone concentrations in the background atmosphere of
668	Southern China, 1994–2007, Atmos. Chem. Phys., 9, 6217–6227, 2009.
669	Wang, T., Xue, L., Brimblecombe, P., Lam, YF., Li, L., and Zhang, L.: Ozone pollution in
670	China: A review of concentrations, meteorological influences, chemical precursors,
671	and effects, Sci. Total Environ., 575, 1582–1596, 2017.
672	Wesely, M. L.: Parameterization of surface resistances to gaseous dry deposition in
673	regional-scale numerical models, Atmos. Environ., 23, 1293–1304, 1989.
674	Xu, J., Tie, X., Gao, W., Lin, Y., and Fu, Q.: Measurement and model analyses of the
675	ozone variation during 2006 to 2015 and its response to emission change in
676	megacity Shanghai, China, Atmos. Chem. Phys., 19, 9017–9035, 2019.
677	Xu, X., Lin, W., Wang, T., Yan, P., Tang, J., Meng, Z., and Wang, Y.: Long-term trend of
678	surface ozone at a regional background station in eastern China 1991–2006:
679	enhanced variability, Atmos. Chem. Phys., 8, 2595–2607, 2008.
680	Xu, W., Lin, W., Xu, X., Tang, J., Huang, J., Wu, H., and Zhang, X.: Long-term trends of
681	surface ozone and its influencing factors at the Mt Waliguan GAW station, China -
682	part 1: overall trends and characteristics, Atmos. Chem. Phys., 16, 6191-6205,
683	https://doi.org/10.5194/acp-16-6191-2016, 2016.
684	Yue, X. and Unger, N.: Ozone vegetation damage effects on gross primary productivity in
685	the United States, Atmos. Chem. Phys., 14, 9137–9153, 2014.





	со	0.31	0.27	0.25	0.23	0.27	0.44	0.56	0.38	0.34
		Ν	NE	Е	SE	S	SW	W	NW	С
689	2012 to	2017.								
688	conditio	ns at Dor	ngtan (DT) site, a r	emote ru	ral site ne	ear the Sh	neshan Is	land (SSI) during
687	southea	st (SE), s	south (S),	southwe	st (SW), v	west (W),	northwes	st (NW) a	nd calm (C) wind
686	Table 1	Mean (CO mixin	g ratios	(ppmv) (under no	rth (N),	northeast	(NE), e	ast (E),





690	Table 2 M	lonthly mean	wind speeds	(m s ⁻¹)	and occurrence	frequencies	(%) of	the
		2		· · · ·			· · ·	

691 southwest (SW) and west (W) winds at Dongtan (DT) site, a remote rural site near the

0	SZ ONESHAN	isianu (001) uu	ining 20	12 10 20	517.							
_		Jan.	Feb.	Mar.	Apr.	May	Jun.	Jul.	Aug.	Sep.	Oct.	Nov.	Dec.
_	SW+W	11.5	9.2	11.9	13.2	12.7	9.8	17.7	10.8	6.1	5.2	11.9	15.1
	Wind speed	2.70	2.93	2.98	3.04	2.86	2.51	2.65	2.77	2.49	2.50	2.55	2.54

692 Sheshan Island (SSI) during 2012 to 2017.





693	Table 3 Statistical results of the Mann-Kendall test and Theil-Sen trend estimate for daily
694	mean values of NO_x , CO mixing ratios, temperature (T), and wind speed (WS) in
695	September and October at Dongtan (DT) site, a remote rural site near Sheshan Island
696	(SSI) site during the 2012–2017 period. The units of the calculated slopes are ppbv yr $^{-1}$ for
697	NO _x and CO, $^\circ\text{C}$ yr^1 for T, and m s^1 yr^1 for WS.

	NO _x	СО	Т	WS
Slope Estimate	0.48*	2.67 [∆]	0.15 [∆]	0.21*

⁶98 ^{*}The result is significant at the 95% confidence level.

 $^{\Delta}$ The result cannot pass the Mann-Kendall trend test at the 90% confidence level.





700	Table 4 Statistical results of the comparisons between the simulated and observed
701	surface O_3 concentrations at Xujiahua (XJH) and Sheshan Island (SSI) sites during
702	September 2014. The calculated O_3 levels are obtained from BC_40, BC_50 and BC_60
703	simulations, respectively. Values of the average surface O_3 concentrations (Mean) and
704	normalized mean bias (NMB) are displayed. The NMB is defined as NMB= $\frac{\sum_{l=1}^{n}(P_{l}-O_{l})}{\sum_{l=1}^{n}O_{l}}$,
705	where P_i and O_i are predicted and observed ozone mixing ratios for sample <i>i</i> , <i>n</i> is the

706 number of total samples (numbers in parentheses).

	Cases	XJH (641)	SSI (720)
	Observation	30.4	57.7
Mean	BC_40	22.0	36.9
(ppbv)	BC_50	25.1	41.8
	BC_60	29.0	47.3
	BC_40	-27.6	-36.1
NMB (%)	BC_50	-17.5	-27.5
	BC_60	-4.6	-18.1





707 Figure Captions

- 708 Figure 1 Land cover of Shanghai and corresponding locations and landscapes of Xujiahui
- 709 (XJH, urban), Dongtan (DT, rural) and Sheshan Island (SSI, remote and oceanic) stations.
- 710 Figure 2 Monthly wind rose diagrams averaged over the period of 2012 to 2017 at
- 711 Dongtan (DT) site, a remote rural site near the Sheshan Island (SSI).
- 712 Figure 3 Monthly mean CO mixing ratios under north (N), northeast (NE), east (E),
- 713 southeast (SE), south (S), southwest (SW), west (W), northwest (NW) and calm (C) wind
- 714 conditions at Dongtan (DT) site, a remote rural site near the Sheshan Island (SSI) during
- 715 **2012 to 2017**.
- 716 **Figure 4** Monthly and year-round mean diurnal variations of O₃ (ppbv) at Sheshan Island
- 717 (SSI, remote and oceanic) and Xujiahui (XJH, urban) sites during 2012 to 2017.
- 718 **Figure 5** Calculated monthly mean ratios of daily maximum O₃ concentrations (O_{3-max}) to

719 minimum O₃ concentrations (O_{3-min}) at Sheshan Island (SSI, remote and oceanic) and

720 Xujiahui (XJH, urban) sites, respectively during 2012 to 2017.

721 Figure 6 Variations of (a) monthly mean O₃ concentrations at Sheshan Island (SSI,

- remote and oceanic) and Xujiahui (XJH, urban) sites during the period 2012-2017, (c)
- corresponding variations of daily mean O₃ concentrations at SSI and XJH in September
 and October, and (c) variations of mean O₃ concentrations during daytime (10:00-16:00
 LST) and nighttime (23:00-04:00 LST) at SSI.
- Figure 7 Daytime and nighttime mean O_3 mixing ratios (ppbv) at Sheshan Island (SSI) and NO_x mixing ratios (ppbv) at Dongtan (DT) site, a remote rural site near SSI under north (N), northeast (NE), east (E), southeast (SE), south (S), southwest (SW), west (W),





- and northwest (NW) wind conditions in MAM (March-May), JJA (June-August), SON
- 730 (September-November), and DJF (December-February), respectively during 2012 to
- 731 **2017**.
- 732 Figure 8 Calculated distributions of monthly mean O₃ concentrations (shades, ppbv) from
- 733 BC_40, BC_50 and BC_60 simulations, respectively in September 2014. Model results
- 734 are compared with observed mean O₃ concentrations (circles, ppbv) obtained from
- 735 Xujiahui (XJH, urban) and Sheshan Island (SSI, remote and oceanic) sites. Also shown is
- 736 the calculated wind field (m s^{-1}) averaged over the same period.
- 737 Figure 9 Mean differences in surface O₃ concentrations (ppbv) simulated with different
- 738 chemical boundaries: (a) BC_50 minus BC_40, (b) BC_60 minus BC_40, and (c) BC_60
- 739 minus BC_50 in September 2014. Also shown is the calculated wind field (m s⁻¹) averaged
- 740 over the simulation period.







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- 742 Figure 1 Land cover of Shanghai and corresponding locations and landscapes of Xujiahui
- 743 (XJH, urban), Dongtan (DT, rural) and Sheshan Island (SSI, remote and oceanic) stations.







- 745 Figure 2 Monthly wind rose diagrams averaged over the period of 2012 to 2017 at
- 746 Dongtan (DT) site, a remote rural site near the Sheshan Island (SSI).







Figure 3 Monthly mean CO mixing ratios under north (N), northeast (NE), east (E),

749 southeast (SE), south (S), southwest (SW), west (W), northwest (NW) and calm (C) wind

750 conditions at Dongtan (DT) site, a remote rural site near the Sheshan Island (SSI) during

751 **2012 to 2017**.







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753 Figure 4 Monthly and year-round mean diurnal variations of O₃ (ppbv) at Sheshan Island

754 (SSI, remote and oceanic) and Xujiahui (XJH, urban) sites during 2012 to 2017.









Jan. Feb. Mar. Apr. May Jun. Jul. Aug. Sep. Oct. Nov. Dec. Jan. Feb. Mar. Apr. May Jun. Jul. Aug. Sep. Oct. Nov. Dec.

756 Figure 5 Calculated monthly mean ratios of daily maximum O₃ concentrations (O_{3-max}) to

757 minimum O₃ concentrations (O_{3-min}) at Sheshan Island (SSI, remote and oceanic) and

758 Xujiahui (XJH, urban) sites, respectively during 2012 to 2017.







corresponding variations of daily mean O₃ concentrations at SSI and XJH in September
and October, and (c) variations of mean O₃ concentrations during daytime (10:00-16:00
LST) and nighttime (23:00-04:00 LST) at SSI.







765

Figure 7 Daytime and nighttime mean O_3 mixing ratios (ppbv) at Sheshan Island (SSI) and NO_x mixing ratios (ppbv) at Dongtan (DT) site, a remote rural site near SSI under north (N), northeast (NE), east (E), southeast (SE), south (S), southwest (SW), west (W), and northwest (NW) wind conditions in MAM (March–May), JJA (June–August), SON (September–November), and DJF (December–February), respectively during 2012 to 2017.







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776	Xujiahui (XJH, urban) and Sheshan Island (SSI, remote and oceanic) sites. Also shown is
777	the calculated wind field (m s ⁻¹) averaged over the same period.







Figure 9 Mean differences in surface O₃ concentrations (ppbv) simulated with different

chemical boundaries: (a) BC_50 minus BC_40, (b) BC_60 minus BC_40, and (c) BC_60

781 minus BC_50 in September 2014. Also shown is the calculated wind field (m s⁻¹) averaged

782 over the simulation period.