1	A measurement and model study on ozone characteristics in marine air at a remote
2	island station and its interaction with urban ozone air quality in Shanghai, China
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11 Abstract

12 To understand the characteristics and changes of baseline ozone (O_3) in oceanic air in 13 East China, a six-year measurement of O₃ concentration was conducted from January 1 14 2012 to September 15 2017 at a remote offshore station located on the Sheshan Island (SSI) near the megacity of Shanghai. The observed monthly mean O₃ concentrations at 15 16 SSI ranged from 33.4 to 61.4 ppbv during the study period, which were about 80% and 12% 17 higher, respectively than those measured at downtown and rural sites in Shanghai. 18 Compared to the remarkable O₃ increases observed at urban and rural sites in Shanghai, 19 observed O₃ concentrations at SSI exhibited statistically insignificant increasing changes (1.12 ppbv yr⁻¹, α >0.10) during the observation period, suggesting less impacts of 20 21 anthropogenic emissions on O₃ levels in oceanic air. In addition, an insignificant 22 decreasing change (-0.72 ppbv yr⁻¹, α >0.10) was detected in O₃ concentrations at SSI in September and October when the influence of regional transport was minimum 23 24 throughout the year, providing a good proxy to study the baseline oxidation capacity of the 25 oceanic atmosphere. City plumes from Shanghai usually carried higher levels of NO_x, 26 resulting in decreased O₃ concentrations at SSI during southwesterly and westerly winds. 27 However, In MAM (March-May) and JJA (June-August), due to the enhanced production of oxygenated volatile organic compounds, O₃ could be continuously produced during 28 29 daytime in aged city plumes, resulting in elevated O_3 concentrations transported to SSI. 30 The impacts of the offshore O₃ on O₃ levels in Shanghai are quantified during an easterly 31 wind dominant episode (September 1-30, 2014) using the WRF-Chem model. Sensitivity 32 results suggest that O_3 in the oceanic air inflows can lead to 20–30% increases in urban

- O_3 concentrations, which should be crucially considered in dealing with urban O_3 pollution
- in large coastal cities like Shanghai.

35 1 Introduction

36 Ground-level ozone (O_3) is a harmful photochemical oxidant detrimental to air quality, 37 human health and land ecosystems (Yue and Unger 2014; Monks et al., 2015; Li et al., 38 2019a). High ambient O₃ has been proved to increase the risks of respiratory and 39 cardiovascular mortality (Goodman et al., 2015) and enhance the greenhouse effect 40 (IPCC, 2013). In recent years, O₃ pollution has drawn increasing attention in China, since 41 O₃ pollution is getting worse in spite of the implementation of Chinese Clean Air Action 42 Plan. In contrast to the 28-40% decreases in $PM_{2.5}$ (fine particulate matter; diameter ≤ 2.5 43 μ m) levels, the observed daily maximum 8-h average (MDA8) O₃ concentrations show increasing rates of 1–3 ppb yr⁻¹ in summer in megacities over eastern China during 2013– 44 45 2017 (Li et al., 2019b). To address the underlying causes of the increasing O₃ pollution 46 has become an urgent issue that triggers lots of discussions based on observational and 47 model studies worldwide (Yang et al., 2014; Lou et al., 2015; Fu et al., 2019). 48 Observational and model studies indicated that the elevated O₃ levels in urban and 49 rural areas in eastern China were strongly related to the changes in anthropogenic 50 emissions of O₃ precursors (Ma et al., 2016; Lu et al., 2018; Li et al., 2019b; Gu et al., 51 2020). Since the O₃ formation was reported to be under volatile organic compound (VOC)

52 limited regime in most Chinese megacities (e.g. Beijing, Shanghai, and Guangzhou), the 53 sharp decreases in nitrogen oxides ($NO_x=NO+NO_2$) emissions combined with slight 54 increases in VOC levels were suggested to be main causes of the observed enhancement 55 of O₃ concentrations in East China (Gao et al., 2017; Xu et al., 2019). In remote areas, 56 changes of baseline O₃ also exhibit sensitive responses to human activities (Vingarzan,

2004; Meng et al., 2009; Wang et al., 2009; Lin et al., 2015). Based on 14-year 57 58 observations at a coastal site in Hong Kong, Wang et al. (2009) pointed out that enhanced 59 pollution flow from the upwind coastal regions contributed to most of the observed O_3 60 increases in the background atmosphere of South China during 1994-2007. And the 61 increase in background O₃, in turn, made a strong contribution of 81% to the increasing 62 rate of O₃ in urban Hong Kong. It is thus necessary to understand the background O₃ 63 changes and their responses to different sources when developing long-term strategies to 64 mitigate local O₃ pollution. However, compared to the intensive field studies in polluted 65 cities and surrounding rural regions, continuous observations of O3 at representative 66 background sites in China are relatively limited (Wang et al., 2017).

67 To better understand the characteristics of the background O₃ changes in mainland 68 China, the China Meteorological Administration (CMA) started to conduct continuous 69 measurements of surface O_3 at several regional background stations (e.g. Shangdianzi, 70 Linan, and Longfengshan) since 2005. Over 10-year records from those sites and 71 Waliguan, a baseline Global Atmospheric Watch (GAW) station in Tibetan Plateau region, 72 exhibited different increases in background continental O₃ concentrations especially 73 during daytime in China (Lin et al., 2008; Xu et al., 2008; Meng et al., 2009; Ma et al., 2016; Xu et al., 2016). The detected positive trends of O₃ were in a range of 0.24–1.13 74 ppbv yr⁻¹, suggesting enhanced atmospheric oxidation capacity of continental air 75 responding to the rapid development of urbanization and industrialization in the past 76 77 decades. In addition to the changes at background O₃ in terrestrial stations mentioned 78 above, the characteristics of baseline O₃ at remote marine sites are also important. It is

79 because that large amounts of O₃ pollution events occurred in coastal urban 80 agglomerations in East China (Lu et al., 2018; Li et al., 2019a, b), affected by both city 81 plumes and oceanic air inflows (Tie et al., 2009; Shan et al., 2016). For example, model 82 work of Tie et al. (2009) suggested that sea air masses carried by oceanic inshore air 83 flows aggravated urban O₃ pollution in Shanghai under convergence conditions. 84 Understanding the O_3 characteristics in offshore oceanic regions is therefore an important prerequisite for understanding the land-sea O₃ interactions and its impacts on O₃ pollution 85 86 in coastal cities. However, to our knowledge, studies on the characteristics and changes 87 of O₃ in marine air are quite limited in mainland China since it is very difficult to conduct 88 systematic and continuous observations under remote oceanic air conditions.

89 In this report, we present the first relatively long and continuous measurements of O_3 90 conducted on a remote offshore island (Sheshan Island, SSI) from January 2012 to September 2017 in eastern China. The SSI is located at the confluence of the Yellow Sea 91 and the East China Sea, covering an area of about 0.4 km². Since there are no inhabitants 92 93 in the island, the observed O_3 is seldom affected by local anthropogenic emissions. The 94 collected O_3 data are used to understand the levels and variabilities of O_3 in the offshore 95 regions and their impacts on the O₃ concentrations in coastal city areas. First shown are 96 the general impacts of regional transport on the remote atmosphere over the SSI region. 97 Then the diurnal patterns of O₃ at SSI are investigated by comparing them with those observed at a downtown site (XJH) in Shanghai. Multi-year changes of O₃ concentrations 98 99 at SSI are analyzed to examine the overall changes of baseline O3 in marine air and 100 possible causes. Also analyzed are the impacts of urban plumes on O_3 levels in oceanic air in offshore regions. At last, the influence of O_3 carried by oceanic air inflows on urban 0₃ air quality in Shanghai is assessed using the Weather Research and Forecasting model coupled with Chemistry (WRF-Chem).

104 2 Material and methods

105 2.1 The SSI site and ozone observations

To investigate the characteristics and variabilities of O_3 in marine air and their interactions 106 with urban air quality in coastal areas, ground O₃ concentrations were continuously 107 108 measured at SSI site (31.4°N, 122.3°E, 73.5 m a.s.l.), which is approximately 75 109 kilometers away from the east edge of Shanghai city. Figure 1 shows the location of SSI 110 and the surrounding environment. As mentioned in Sect. 1, there is no resident and tourist on the island. The observed O₃ at SSI site can represent the background O₃ conditions in 111 112 oceanic air which are seldom contaminated by anthropogenic emissions. Hourly O₃ data 113 was collected during January 1 2012 to September 15 2017, with a capture rate of 89.7%. O₃ was measured using an analyzer from Ecotech, Australia (Model EC9810), which 114 115 combined microprocessor control with ultraviolet photometry. The instrument met the 116 technical specifications for United States Environmental Protection Agency, with a quality control check every 3 days, filer replaced every 2 weeks and calibration every month. 117

118 2.2 Observational data at urban and rural sites in Shanghai

To better understand the characteristics of the offshore O_3 in oceanic air at SSI, O_3 observations obtained from a downtown site, Xujiahui (XJH) are used for comparisons. The XJH site is located at downtown Shanghai, approximately 80 km west from the SSI. Since measurements of NO_{x_1} carbon monoxide (CO) and meteorological parameters (e.g. 123 wind direction and wind speed) were unavailable at SSI, observations obtained at an adjacent site, Dongtan (DT), are substituted for the investigation. The DT site was set up 124 in a national nature reserve near the coast of Shanghai, where the observed pollutant 125 levels have been reported to well reflect the impacts of megacities in the Yangtze River 126 127 Delta (YRD) region on the remote atmosphere during the MIRAGE-Shanghai (Megacities Impact on Regional and Global Environment at Shanghai) field campaign (Tie et al., 2013). 128 Similar to SSI, the DT site is also little affected by human activities. The obtained 129 130 observations of meteorology and pollutants are therefore applied for analyzing the 131 impacts of regional transport on observed O₃ concentrations at SSI. NO_x concentrations 132 were measured with a chemiluminescent trace level analyzer (TEI; Model 42iTL), with detection limit of 0.025 ppb. CO concentrations were measured by the Model 48iTL 133 134 Enhanced CO analyzer, based on gas filter correlation technology. The wind speed and 135 wind direction were measured by using a DZZ4 Automatic Weather Station certificated by the China Meteorological Administration. The geographical locations and surrounding 136 137 environment of XJH, DT, and SSI are displayed in Fig. 1.

138 2.3 The WRF-Chem model

We simulate O₃ using the regional chemical transport model WRF-Chem (version 3.8, https://www2.acom.ucar.edu/wrf-chem), collaboratively developed through efforts of several institutes, such as the National Center for Atmospheric Research (NCAR) and the National Oceanic and Atmospheric Administration (the National Centers for Environmental Prediction (NCEP). The model includes on-line calculation of meteorological parameters, transport, mixing, emission, and chemical transformation of trace gases and aerosols

(Grell et al., 2005). The Regional Acid Deposition Model version 2 (RADM2, Stockwell et 145 al., 1990) gas-phase chemical mechanism is used for the O₃ formation chemistry. 146 147 Photolysis rates are calculated by using the fast radiation transfer module (FTUV) followed those in Madronich and Flocke (1999) and Tie et al. (2003). ISORROPIA II 148 149 secondary inorganic (Fountoukis and Nenes, 2007) and the Secondary ORGanic Aerosol 150 Model (SORGAM) (Schell et al., 2001) schemes are used for aerosol chemistry. Dry 151 deposition follows the standard resistance-in-series model of Wesely (1989). The major 152 physical processes employed in the model follow the Lin microphysics scheme (Lin et al., 1983), the Yonsei University (YSU) planetary boundary layer (PBL) scheme (Hong and 153 Lim, 2006), the Noah Land surface model (Chen and Dudhia, 2001), and the long-wave 154 155 radiation parameterization (Dudhia, 1989).

156 The model used in this study has a horizontal resolution of 6km×6km, including 150 157 un-staggered grids in west-east, 150 un-staggered grids in south-north, and 35 vertical layers extending from the surface to 50 hPa. The domain encompasses Shanghai and its 158 159 surrounding region, centered at 31.3°N, 121.4°E. The NCEP FNL (Final) Operational 160 Global Analysis data are used for meteorological initial and boundary conditions, with lateral meteorological boundary updated every 6 h. Basic chemical lateral boundary 161 conditions are constrained by a global chemical transport model (MOZART-4, Model for 162 163 OZone And Related chemical Tracers, version 4) (Tie et al., 2001; Emmons et al., 2010). Anthropogenic emissions are derived from the Multi-resolution Emission Inventory for 164 165 China (MEIC inventory, http://www.meicmodel.org/; Li et al., 2014) for year 2010. Biogenic 166 emissions are calculated online using model of emissions of gases and aerosols from

167 nature (MEGAN2, Guenther et al., 2006).

168 2.4 Methods for assessing the trend of ozone

169 The daily mean O₃ concentrations are used to examine the overall changes in O₃ concentrations during the period 2012-2017, including all time of day with qualified 170 171 measurements. The trends are assessed using two nonparametric methods, which are 172 commonly used to detect trends of non-normally distributed data with seasonality (Xu et al., 2016).The Mann-Kendall (MK) trend test (Mann, 1945; Kendall, 1975; Gilbert, 1987) is 173 174 used to examine the trend significance, and the Theil-Sen trend estimate method (Sen, 175 1968) is used to estimate the slope of trend, which could also be considered as the rate of 176 change, during the six-year period. Compared to the linear fitting analysis which requires data to be independent and follow a Gaussian distribution, the non-parametric trend test 177 178 methods only need the data to be independent (Gocic and Trajkovic, 2013). To determine 179 if the calculated rate of change is statistically significant, the confidence level of at least 95% 180 is adopted in the MK trend test, with α value less than 0.05 being considered a statistically 181 significant trend. The trend significance is examined by comparing the value of a 182 standardized test statistic Z to that of a standard normal variate at a given significance level (Z_{α} , α =0.05). If $|Z| > Z_{1-\alpha/2}$, then the dataset is non-stationary, exhibiting either an 183 184 increasing or a declining trend; If $|Z| \leq Z_{1-\alpha/2}$, then the dataset is stationary with no significant trend. Detailed calculation of Z can be referred to Xu et al. (2016). 185

186 **3 Results and discussion**

187 **3.1** Regional transport characteristics at SSI

188 The observed O₃ concentrations at SSI were inevitably influenced by regional transport

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depending on the prevailing winds in various seasons. Figure 2 displays the monthly wind 189 190 rose diagrams averaged over the period of 2012 to 2017 at DT. As mentioned in Sect. 2.2, the DT site is a rural site located quite close to SSI. The observed wind speeds and wind 191 192 directions could then be applied to deduce the origins of the air mass arriving at SSI in 193 adjacent region. Generally, observed prevailing winds exhibited distinct seasonal 194 variabilities which were greatly affected by the East Asian monsoon. In warm seasons (May-August), the site was predominately influenced by easterly and southeasterly winds, 195 196 accounting for 40–50% of the total winds. While in cold seasons (November-February), 197 the northwesterly and northerly winds became the predominant flows that affected SSI, 198 accounting for about 45% of total winds. During transitional months (e.g. March, April, 199 September and October), the dominant winds presented more diversities, with wind 200 directions dispersedly distributed in all the directions. The observed seasonal variations of 201 prevailing winds are typical at coastal cities at mid-latitude region (Shan et al., 2016; Xu et 202 al., 2019), suggesting that air masses arriving at SSI originated from various regions and 203 could result in different impacts on the offshore atmospheric composition in different 204 months.

Since CO has a relative long chemical lifetime of a few months, the observed CO concentrations at DT could be regarded as a consequence of regional transport from polluted regions (Tie et al., 2009). Figure 3 displays the observed monthly mean CO mixing ratios under wind directions of north (N), northeast (NE), east (E), southeast (SE), south (S), southwest (SW), west (W), and northwest (NW) at DT during the 2012–2017 period. Observed CO exhibited relative higher concentrations under SW and W winds in

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all months, with mean mixing ratios of 0.44 and 0.56 ppmv, respectively during 2012-2017 211 212 (Table 1). The observed high CO mixing ratios suggested that the atmosphere 213 constituents at SSI could be more affected by regional transport of air pollutants under 214 SW and W wind conditions. As SSI is located to the northeast of the Shanghai city (Fig. 1), 215 air masses carried by the SW and W flows usually contain more urban pollutants from 216 upwind city areas, and those carried by E, SE, and NE flows mostly come from the ocean. 217 The oceanic air masses are less polluted compared to those from the cities, leading to 218 lower CO mixing ratio at SSI. For example, observed CO exhibited a mean concentration 219 of 0.23 ppmv under SE wind conditions, which was about 50% lower than that influenced 220 by W winds. To further examine the impacts of the SW and W winds on the atmosphere constituents at SSI, Table 2 lists the calculated monthly mean occurrence frequency of the 221 222 SW and W winds in separate months during the studied period. The SW and S winds 223 were most infrequent in September (6.1 %) and October (5.2 %), suggesting that the 224 atmosphere at SSI during the two months could be less contaminated by pollutants 225 transported from the city and might be more close to the baseline oceanic air conditions.

3.2 The diurnal pattern of ozone at SSI

Figure 4 displays the monthly mean diurnal variations of O₃ at SSI and XJH in different months during 2012–2017. The observed O₃ concentrations at the two sites exhibited similar seasonal variations, with monthly mean values highest (61.4 ppbv for SSI and 35.9 ppbv for XJH) in May and lowest (33.4 ppbv for SSI and 12.5 ppbv for XJH) in December. Since the O₃ formation in urban Shanghai is VOC-limited, observed O₃ could be significantly depressed by large NO_x emissions at downtown site (XJH) (Gu et al., 2020). 233 In Fig. 4, observed O_3 levels at XJH were quite lower than those at SSI in all months, with 234 mean concentrations of 27.8 and 50.1 ppbv, respectively at XJH and SSI during the 235 observation period. The observed mean daily maximum 8-h average (MAD8) O₃ concentrations exhibited same differences between the two sites, which were 40.1 and 236 62.0 ppbv, respectively at XJH and SSI. The observed mean O₃ concentration at SSI was 237 238 also higher than that at DT (44.7 ppbv, Fig. S1) which is more close to the city, suggesting 239 that O_3 levels in marine air could be higher than those at continental urban and rural sites. 240 The observed diurnal patterns of O₃ at SSI and XJH in Fig. 4 were similar to those 241 reported for other sites in eastern China (Xu et al., 2008; Geng et al., 2015; Gao et al., 2017), exhibiting minimums in early morning (06:00-08:00 LST) and maximums in the 242 243 afternoon (13:00-15:00 LST). However, compared to those at the urban site (XJH), 244 observed amplitudes of O₃ diurnal variations were much smaller at SSI. The diurnal variations of surface O₃ can be mainly attributed to the O₃ production through 245 246 photochemical reactions in the daytime and O_3 depression via NO titration at nighttime 247 (Sillman, 2003). Due to few emissions of O_3 precursors (NO_x and VOCs), the O_3 248 production and depression could be weaker at remote site, resulting in flatter diurnal cycle of O₃ compared to that at polluted urban site. 249

Since the amplitudes of O_3 diurnal variations usually exhibited much smaller values in background areas compared to those in polluted urban regions, the ratio of daily maximum O_3 concentration (O_{3-max}) to minimum O_3 concentration (O_{3-min}) was regarded as an indicator to identify if the local O_3 pollution was significantly influenced by anthropogenic emissions (Cvitas and Klasinc 1993; Vingarzan, 2004). The O_{3-max}/O_{3-min}

ratio displayed larger values in polluted regions (Cvitas et al., 1995) and lower values in 255 256 less contaminated rural regions. A ratio of about 1.4 suggested that the site could be 257 regarded as a typical background site (Scheel et al., 1997). For regional background sites in China, the typical values of O_{3-max}/O_{3-min} were usually in the range of 2–3 (Xu et al., 258 259 2008; Meng et al., 2009; Gu et al., 2020). In Lin'an, a continental background site in YRD 260 region, the ratio was reported to increase as a result of NO_x emission changes during past decades, which could reach above 6 during summertime (Xu et al., 2008). In Fig. 4, 261 observed O₃ displayed different diurnal variabilities in various months at SSI. The 262 263 variations of the O_{3-max}/O_{3-min} ratio suggested different influence of regional transport on 264 O_3 levels in the marine atmosphere.

265 Figure 5 displays the calculated monthly mean O_{3-max}/O_{3-min} ratios at SSI and XJH, respectively during 2012–2017. Generally, the observed ratios of O_{3-max}/O_{3-min} at SSI were 266 much lower than those at XJH in all the months, suggesting less impact of anthropogenic 267 emissions on O_3 levels. The calculated mean ratios were 3.03 and 5.20, respectively at 268 269 SSI and XJH, and most of the calculated values were larger than 4.50 at the urban site. 270 Besides, the ratios presented distinct seasonal differences at XJH and SSI sites. Higher values were observed in summer, indicating stronger photochemical production of 271 daytime O_3 during June to August. At SSI, the O_{3-max}/O_{3-min} ratio exhibited relatively low 272 273 values in September and October, ranging from 1.61-2.35 during the studied period. The values were consistent with the typical values of O_{3-max}/O_{3-min} observed at continental 274 275 background sites in China (Xu et al., 2008; Meng et al., 2009; Gu et al., 2020). Since the 276 observed temperature and solar radiation still exhibited higher values during the two

months in Shanghai (Gao et al., 2017), the observed low O₃ diurnal amplitudes should not 277 278 be attributed to the weakened photochemical formation of O₃ as those in winter. Due to 279 the persistent control of anticyclone, Shanghai and its neighboring areas are usually 280 dominated by stable weather conditions in September and October, resulting in more 281 gentle and diversified wind conditions. During the two months, the occurrences of more 282 polluted SW and W winds were lowest (6.1% and 5.2%) throughout the year. The corresponding wind speed (2.49 and 2.50 m s⁻¹) also exhibited values 20% lower those in 283 284 other months (Table 2). The transport conditions led to fewer pollutants transported to the 285 SSI region, which could explain the observed weak diurnal variabilities of O₃ in September 286 and October. The transport conditions together with O_3 response further confirmed that the transport of city pollutants had minimum impacts on the offshore O₃ levels in oceanic 287 288 air at SSI in September and October, providing a good proxy to study the baseline oceanic 289 O_3 and oxidation capacity of background atmosphere in eastern China. 290 3.3 Overall changes of ozone in oceanic air at SSI 291 Several studies have observed increasing trends of ground-level O_3 in metropolitan areas 292 over eastern China since 2013, suggesting that the O_3 increases were mostly attributed to the NO_x emission reductions (Ma et al., 2016; Gao et al., 2017; Lu et al., 2018; Li et al., 293 294 2019b). However, the O_3 changes at remote sites were relatively not well elucidated 295 during past years. Figure 6a presents the monthly variations of O₃ concentrations at SSI and XJH during the 2012-2017 period. The statistical results of the MK test and Theil-296 297 Sen trend estimate method indicated that observed monthly mean O₃ mixing ratios (O_{3-ave}) 298 exhibited increasing changes at both urban (XJH) and remote sites (SSI) in Shanghai,

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with calculated increasing rate of 1.97 and 1.12 ppbv yr⁻¹, respectively in XJH and SSI. 299 Though an overall upward trend of O₃ was detected at SSI, the changes were not as 300 301 remarkable as those observed at XJH, which could not even pass the MK trend test at the 302 90% confidence level. The monthly mean MDA8 and daily extreme values of O_3 exhibited 303 similar differences between the two sites. The calculated increasing rates of MDA8 O_{3} , O_{3-max} and O_{3-min} were 2.73, 2.77, and 1.35 ppbv yr⁻¹ (α <0.05), respectively at XJH, and 304 1.01, 1.35, and 1.27 ppbv yr⁻¹ (α >0.10), respectively at SSI. Compared to the rapid O₃ 305 306 increases in urban Shanghai which was suggested to mostly result from the significant 307 NO_x emission reductions in the past decade (Gao et al., 2017; Xu et al., 2019), the 308 statistically insignificant changes of O_3 detected at SSI indicated that O_3 in the oceanic air 309 remained a relative constant level during the study period and was less influenced by the 310 decreases of NO_x emissions.

311 As discussed in Sect. 3.1, the prevailing winds carried different levels of pollutants to 312 the SSI, resulting in different impacts on the O₃ levels in different months. In September 313 and October, the frequencies of SW and W winds that carried high levels of pollutants 314 were lowest (Table 1–2), exerting least influence on the atmospheric composition at SSI. 315 Therefore, the variations of surface O₃ concentrations in September and October at SSI were examined to further assess the changes of least contaminated O_3 in the oceanic air. 316 Figure 6b presents the overall changes of daily mean surface O3 concentrations in 317 September and October at SSI and XJH, respectively during the six-year period. The 318 319 corresponding mean O₃ mixing ratios during the two months were 60.9 and 31.3 ppbv, 320 respectively at SSI and XJH. Compared to the significant elevated O_3 concentrations at 321 XJH (0.59 ppbv yr⁻¹, α <0.10) in September and October, observed O₃ at SSI during same 322 months exhibited insignificant decreasing changes from 2017–2017. The changes (-0.72 323 ppbv yr⁻¹, α >0.10) were somewhat different from the overall O₃ changes (+1.12 ppbv yr⁻¹, 324 α >0.10) at SSI, suggesting different causes of the observed O₃ changes in the oceanic air 325 during September and October.

326 To investigate possible drivers of the observed changes in the least contaminated O_3 327 in September and October at SSI, Table 3 displays the statistical results of the MK test 328 and Theil-Sen trend estimate for NO_x and CO mixing ratios, temperature, and wind speed 329 during the 2012–2017 period. Statistically significant upward trends were detected in wind speed, with estimated increasing rates of 0.21 m s⁻¹ yr⁻¹ during the observation period 330 (α <0.05). The significantly enhanced surface wind speeds were conducive to the diffusion 331 332 of O₃, which might be an important meteorological driver of the observed decreasing changes in O₃ levels at SSI from 2012 to 2017. Observed NO_x and CO levels exhibited 333 increases of 0.48 ppbv yr⁻¹ (α <0.05) and 2.67 ppbv yr⁻¹ (α >0.10), respectively in 334 335 September and October during the six-year period, indicating enhanced transport of 336 pollutants to the oceanic area. Tie et al. (2013) suggested that the VOC-limited regime of O₃ formation was not only confined in urban Shanghai, but also extended to a broader 337 338 regional area surrounding Shanghai. Thus, the elevated NO_x concentrations might not 339 only retard daytime O₃ production but also enhance nighttime O₃ depression at SSI. Figure 6c further presents corresponding variations of daytime (10:00-16:00 LST) and 340 341 nighttime (23:00-04:00 LST) mean O₃ concentrations at SSI. Both daytime and nighttime O₃ concentrations exhibited downward changes, reflecting the O₃ response to the 342

enhanced O_3 diffusion and depression in September and October. Therefore, the enhanced diffusion and depression of O_3 induced by the elevated wind speed and NO_x concentrations might be important causes of the observed O_3 changes in September and October at SSI. It should be noted that the influence of radiation cannot be analyzed since observations of solar radiation were not available during the study period. Therefore, more measurements are still needed to further understand the O_3 changes and corresponding drivers in the oceanic air.

350 3.4 Impacts of urban plumes on ozone in oceanic air at SSI

351 Due to the relatively long residence lifetime (about one month), O_3 produced at urban 352 regions could be transported several hundred kilometers away to downwind areas. Meanwhile, the urban plumes become more aged with continuous production/depletion of 353 354 O₃ and its precursors, resulting in non-linear changes in O₃ in downwind areas (Geng et 355 al., 2011; Tie et al., 2009, 2013). Several studies suggested that there tended to be considerable O₃ formations in aged urban plumes in the downwind region of Shanghai 356 357 (Geng et al., 2011; Tie et al., 2013). To investigate the impacts of urban plumes on the O_3 358 levels in oceanic air at SSI, the relationships between observed O_3 and NO_x under different wind conditions at SSI and DT are investigated in this section. 359

Figure 7 presents the daytime and nighttime O_3/NO_x -wind relationships in MAM (March–May), JJA (June–August), SON (September–November), and DJF (December– February), respectively during 2012–2017. The SW and W winds were associated with higher NO_x concentrations in both daytime and nighttime. The result was consistent with the observed CO changes in Sect. 3.1. Since there is no local anthropogenic emission at

SSI, the higher levels of NO_x and CO were mainly resulted from the transport of more 365 polluted urban plumes by the SW and W winds. Generally, observed daytime O3 and NOx 366 367 concentrations presented opposite variations with the wind direction changes (Fig. 7a). In SON and DJF, the correlation coefficients (Rs) between daytime O_3 and NO_x were -0.72 368 and -0.75, respectively, indicating that the O₃ formation was inhibited by increased NO_x 369 370 concentrations. The results are in accordance with Tie et al. (2013) and Xu et al. (2019), 371 who suggested that Shanghai and a broader regional area surrounding the city were all in 372 the VOC-limited O_3 formation regime during the study period. However, in MAM and JJA, 373 the daytime O₃-NO_x variations presented totally different patterns under SW and W wind conditions. As wind directions turned from E-SE to SW-W, observed mean NO_x 374 375 concentrations increased from about 10 ppbv to 20 ppbv, while observed mean O₃ 376 concentrations increased from 50-60 ppbv to 70-80 ppbv. The enhancements in daytime O₃ levels suggested that there should be persistent production of O₃ in the polluted air 377 378 masses carried by the SW and W winds in MAM and JJA.

379 Based on observations and WRF-Chem simulations, Tie et al. (2013) suggested 380 considerable O₃ production in aged city plumes in the downwind area of Shanghai. Since air masses affecting SSI site were directly originated from Shanghai under the SW and W 381 382 wind conditions (Fig. 1), the observed O_3 enhancements should be mainly attributed to the 383 O₃ production in the city plumes carried by SW and W winds. Studies during the MIRAGE-Shanghai campaign suggested several factors that contributed to the O3 384 385 enhancements in aged city plumes downwind Shanghai. First, as there is a large area of forest located in the south of Shanghai, Geng et al. (2011) suggested that continuous 386

387 oxidation of isoprene emitted by the biogenic sources could result in enhanced production of hydrogen radicals (HO₂) especially in warm seasons. Once the air massed were 388 389 transported north and mixed with high NO_x emissions, O₃ would be quickly produced. 390 However, the impacts of biogenic emissions on O_3 production were mainly limited in the 391 south part of Shanghai, which can hardly influence the atmosphere in the SSI region. 392 Then, Tie et al. (2013) further illustrated that the OH reactivity of alkane, alkene, aromatics, and oxygenated VOCs (OVOCs) contributed to the O3 formation in city plumes. Among 393 394 them, the influence of alkane, alkene and aromatics mostly occurred within or near the city, 395 while the OVOCs could be produced or emitted during the transport of the city plumes, 396 resulting in substantial O_3 enhancements in aged city plumes at 100–200 km downwind 397 Shanghai.

398 The SSI is located approximately 100 km northeast from the downtown area of Shanghai. In MAM and JJA, the SW and W winds carried air masses with enhanced 399 400 OVOCs oxidation and O_3 production, resulting in elevated daytime O_3 levels on the island. 401 While in SON and DJF, the observed O₃ decreases at SSI during SW and W winds 402 suggested lower efficiency of O_3 productivity in the city plumes. That might be because that fewer OVOCs were released or produced downwind the city due to the lower 403 temperature and weaker solar radiation (Cai et al., 2009). In addition, in SON and DJF, the 404 405 SW and W winds were usually related to low pressure system with large cloud cover and rich water vapor in Shanghai, which could also lead to depressed photochemical reactions 406 407 and decreased O₃ levels. At night, observed O₃ and NO_x displayed totally opposite 408 changes with wind directions (Fig. 7b), indicating O_3 depression by nighttime NO_x titration

in all the seasons. High O_3 levels were observed under northeasterly, easterly and southeasterly oceanic wind conditions, ranging from 50–60, 30-55, 55–60, and 40–50

412 3.5 Impacts of offshore ozone on urban ozone air quality in Shanghai

ppbv respectively at night in MAM, JJA, SON, and DJF.

411

413 As is presented in Sect. 3.2 and 3.3, observed O₃ concentrations at SSI were much higher than those at urban site (XJH), suggesting higher levels of O_3 in oceanic air than those on 414 the continent. Therefore, sea breezes tend to bring more O3 to the continent, aggravating 415 416 O₃ pollution in coastal cities. Shanghai is one of the largest cities located on the east coast 417 of China, experiencing severe O₃ pollution in recent years (Xu et al., 2019; Gu et al., 418 2020). According to the cluster analysis results (Fig. S2), easterly winds from the ocean greatly affected the Shanghai region, accounting for 64-78% of the total flows in 419 420 non-winter months during the period 2012–2017. To understand the impacts of higher O₃ 421 in oceanic air on the urban air quality, numerical experiments are conducted using the 422 WRF-Chem model to examine the response of O_3 levels in Shanghai to various oceanic 423 air inflow conditions in this section.

Simulations are performed during September 1–30 2014 when the prevailing winds were mostly northeasterly and easterly in the Shanghai region. The occurrence frequencies of the northeasterly and easterly winds were 23% and 27% respectively, during the simulation period, suggesting dominant influence of the oceanic air inflows on the city of Shanghai. Consistent with above analysis, observed air O₃ concentrations were much higher in oceanic regions than those in city areas, with monthly mean values of 30.9 and 57.7 ppby, respectively at XJH and SSI in September 2014. The chemical boundary

21

431 conditions (BCs) of the regional model can represent the inflows conditions to explore 432 their impacts on surface concentrations of air pollutants over a certain continent region. 433 Using this method, Pfiter et al. (2011) proposed that chemical inflows taken from different observational and model datasets could result in differences of ± 15 ppbv in O₃ levels in 434 435 the US west coast region. Therefore, three sets of numerical experiments are conducted as follows to access the impacts of oceanic O₃ air inflows on the urban O₃ air quality in 436 Shanghai. All the simulations are driven by the same emissions, initial conditions, physical 437 438 and chemical schemes.

439 (1) BC_40: O_3 concentrations at the eastern lateral boundary of the domain on the 440 ocean are assigned to 40 ppbv, which is provided by the MOZART-4 model, closed to the 441 observed urban O_3 levels (29.0–38.4 ppbv) in Shanghai in September. The chemical BCs 442 are updated every 6 hours.

443 (2) BC_50: Same as BC_40, but with O₃ concentrations setting to 50 ppbv at the
444 eastern lateral boundary of the domain.

445 (3) BC_60: Same as BC_40, but with O_3 concentrations at the eastern lateral 446 boundary of the domain setting to 60 ppbv according to the observed O_3 levels at SSI 447 (50.9–71.0 ppbv) in September.

Figure 8 displays the simulated and observed monthly mean distributions of surface O₃ concentrations in BC_40, BC_50 and BC_60 scenarios, respectively. In addition to the observations at XJH and SSI, O₃ measurements obtained from other three sites, Pudong (PD, suburban), Sheshan (SS, rural), and Dongtan (DT, rural), during the same period were introduced to evaluate the model's performance in simulating O₃ in Shanghai. The O₃ concentrations at all the sites were measured using the same method as described in Sect. 2.1. The calculated distributions of O₃ agree with observations, which exhibit lower values in urban regions compared to those in rural and ocean areas, indicating strong O₃ depressions in the city of Shanghai due to the VOC-limited O₃ formation regime. The R values between the simulated and observed O₃ concentrations are all larger than 0.50 at continental sites (XJH, PD, SS, and DT), suggesting good prediction of O₃ variations by the model.

460 Table 4 displays the statistical results of the comparisons between the simulated and observed surface O₃ concentrations at different sites in Shanghai. Generally, the 461 462 WRF-Chem model underestimates O₃ concentrations at all the sites in most cases. Taken the BC_40 scenario for example, the O₃ concentrations are underestimated by 9.4–27.6% 463 464 at continental sites and 36.1% at SSI, suggesting larger underestimation of O3 concentrations in oceanic regions. Model results further suggest that elevated O_3 levels in 465 the eastern chemical BCs would lead to increases in O_3 concentrations at both urban and 466 467 remote sites when the prevailing winds are mostly easterly in Shanghai. With O_3 468 concentrations increasing from 40 to 60 ppbv in the easterly oceanic air inflows, the simulated monthly mean O₃ concentrations increase by 7.0–9.7 ppbv at continental sites 469 and 10.4 ppbv at SSI. The underestimation of O₃ levels by the model is also greatly 470 471 improved in the BC_60 scenario, when the chemical BCs of O3 are more close to the observations. Compared to those in the BC_40 scenario, the normalized mean bias 472 473 (NMBs) of the predicted O₃ concentrations reduced at most sites in the BC_60 scenario, for example from -36.1 % to -18.1 % at SSI and -27.6% to -4.6% at XJH, suggesting a 474

475 crucial role of the eastern oceanic air inflows in influencing O₃ air quality in Shanghai.

The calculated monthly mean differences in surface O₃ concentrations between 476 477 simulations in different scenarios are further presented in Fig. 9. Since the dominant winds are easterly during the simulation period, distinct changes in surface O_3 concentrations 478 479 throughout Shanghai are generated, exhibiting generally gradient increases from the ocean to the continent as O₃ increases in the oceanic air inflows. With every 10 ppbv 480 increases in O₃ levels in oceanic air, the simulated surface mean O₃ concentrations 481 482 increase by 3-6 ppbv in the land area and 4-7 ppbv in the offshore region. Due to the 483 strong O₃ depressions associated with high anthropogenic emissions, the simulated O₃ 484 enhancements are relatively lower in the central urban region compared to those in surrounding areas. Even so, simulated mean O₃ concentrations still exhibit 6–8 ppbv 485 486 increases in downtown Shanghai in the BC_60 scenario, accounting for approximately 30% of the simulated O₃ concentrations in the BC_40 case. During the period 2012–2017, 487 most of the measured O_3 concentrations ranged between 50–60 ppbv at SSI in non-winter 488 489 seasons. Carried by the easterly inflows, these oceanic air masses with higher O_3 levels 490 (50-60 ppbv) could be transported to the coastal regions, resulting in approximately 20-30% increases in urban O₃ concentrations in Shanghai according to the sensitivity results. 491 492 4 Conclusions 493 In this paper, we present the first relatively long and continuous measurements of oceanic air O₃ conducted at an offshore monitoring station on the Sheshan Island during January 1 494 495 2012 to September 15 2017. The southwesterly and westerly winds are proved to carry

496 more pollutants to the SSI site, exerting greater influence of human activities on the

497 oceanic atmosphere over the offshore region of the East China Sea. Since the two kinds
498 of winds exhibited minimum occurrence frequencies and wind speeds in September and
499 October, atmosphere at SSI during the two months are considered to be less affected by
500 the transport of regional pollution.

501 Compared to those in urban (XJH) and rural (DT) sites, the observed O₃ levels were 502 higher at SSI, with a mean value of 50.1 ppbv during the observation period. Similar seasonal and diurnal patterns of O3 were observed at SSI and XJH; however, the 503 amplitudes of O₃ variations were much smaller at the offshore site (SSI). Since O₃ 504 505 formation in Shanghai and its surrounding regions were VOC-limited, the observational results suggested that the production and depression of O₃ could be weaker in the ocean 506 regions due to weak influence of the anthropogenic emissions. Observed mean 507 508 O_{3-max}/O_{3-min} ratios also exhibited lower values at SSI (3.03) than those at XJH (5.20), with minimum values ranging from 1.61-2.35 in September and October. The result further 509 510 illustrated that SSI was seldom affected by the anthropogenic emissions, especially in 511 September and October.

The multi-year changes of the oceanic O_3 at SSI are investigated using the Mann-Kendall trend test and the Theil-Sen trend estimate method during 2012–2017. Different from the significant O_3 increases detected at XJH and other rural sites reported in previous studies, the observed mean O_3 concentrations at SSI exhibited statistically insignificant increasing changes (1.12 ppbv yr⁻¹, α >0.10) during the observation period and insignificant decreasing changes (-0.72 ppbv yr⁻¹, α >0.10) in September and October when the transport of city pollutants had minimum impacts on the island. Due to fewer impacts of anthropogenic emissions, most of the observed changes in O_3 at SSI could be attributed to the changes of meteorological conditions. Observed wind speed exhibited significant increases (0.21 m s⁻¹ yr⁻¹, α <0.05) in September and October during the observation period, suggesting that enhanced diffusion conditions could be an important meteorological factor in determining the decreases in O_3 concentrations during the observation period.

The impacts of urban plumes on O₃ levels in oceanic air at SSI are evaluated by 525 526 studying the relationships between observed O_3 and NO_x under different wind conditions. 527 The SW and W winds usually carried air masses with higher NO_x concentrations in both daytime and nighttime to the island. Generally, observed daytime and nighttime O₃ 528 529 concentration decreased as NO_x concentration increases in SW and W winds, exhibiting 530 typical VOC-limited characteristics of O₃ formation. The pattern was more typical in SON and DJF, with R values of -0.72 and -0.75, respectively between O3 and NOx 531 532 concentrations. In MAM and JJA, the daytime O₃-NO_x variations presented kind of 533 positive relationships under SW and W wind conditions, suggesting continuous O_3 534 production in aged city plumes from Shanghai. As reported in previous studies during the MIRAGE-Shanghai campaign, enhanced OVOCs oxidation should be the most important 535 536 driver of the observed O₃ enhancements in the city plumes transported by the SW and W 537 winds.

The influence of the oceanic O_3 air inflows on urban O_3 air quality in Shanghai are quantified during an easterly wind dominant episode (September 1–30, 2014). Numerical experiments are conducted with chemical BCs of O_3 assigned according to different inflow

541	conditions using the WRF-Chem model. Model results suggest that increases of O_3 in the
542	easterly oceanic air inflows will lead to gradient increases from the ocean to the continent.
543	With every 10 ppbv O_3 increases, the calculated surface mean O_3 concentrations can
544	increase by 3–6 ppbv in the land and 4–7 ppbv in the offshore region. Compared to those
545	in surrounding regions, O_3 in central city of Shanghai exhibited lower enhancements in
546	response to the O_3 increases in oceanic air inflows due to strong O_3 depression processes.
547	Even so, the impacts of the oceanic air inflows can still lead to 20–30% increases in urban
548	O_3 concentrations which should be crucially considered in dealing with O_3 pollution in
549	large coastal cities like Shanghai.
550	
551	Data availability. The data used in this paper can be provided upon request from Dr.
552	Jianming Xu (metxujm@163.com).
552 553	Jianming Xu (metxujm@163.com).
	Jianming Xu (metxujm@163.com). Author contribution. YG and JX came up with the original idea, designed the analysis
553	
553 554	Author contribution. YG and JX came up with the original idea, designed the analysis
553 554 555	Author contribution. YG and JX came up with the original idea, designed the analysis methods, developed the model code, and performed the simulations. WG provided the
553 554 555 556	<i>Author contribution.</i> YG and JX came up with the original idea, designed the analysis methods, developed the model code, and performed the simulations. WG provided the observational data. YG and YQ conducted the analysis of the observations and model
553 554 555 556 557	<i>Author contribution.</i> YG and JX came up with the original idea, designed the analysis methods, developed the model code, and performed the simulations. WG provided the observational data. YG and YQ conducted the analysis of the observations and model
553 554 555 556 557 558	<i>Author contribution.</i> YG and JX came up with the original idea, designed the analysis methods, developed the model code, and performed the simulations. WG provided the observational data. YG and YQ conducted the analysis of the observations and model results. YG prepared the manuscript with contributions from all co-authors.
553 554 555 556 557 558 559	<i>Author contribution.</i> YG and JX came up with the original idea, designed the analysis methods, developed the model code, and performed the simulations. WG provided the observational data. YG and YQ conducted the analysis of the observations and model results. YG prepared the manuscript with contributions from all co-authors.

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727	Table 1 Mean CO mixing ratios (ppmv) under north (N), northeast (NE), east (E),
728	southeast (SE), south (S), southwest (SW), west (W), northwest (NW) and calm (C) wind
729	conditions at Dongtan (DT) site, a remote rural site near the Sheshan Island (SSI) during
730	2012 to 2017.

	Ν	NE	Е	SE	S	SW	W	NW	С
СО	0.31	0.27	0.25	0.23	0.27	0.44	0.56	0.38	0.34

Table 2 Monthly mean wind speeds (m s⁻¹) and occurrence frequencies (%) of the 731 southwest (SW) and west (W) winds at Dongtan (DT) site, a remote rural site near the 732 Sheshan Island (SSI) during 2012 to 2017.

733

	Jan.	Feb.	Mar.	Apr.	May	Jun.	Jul.	Aug.	Sep.	Oct.	Nov.	Dec.
SW+W	11.5	9.2	11.9	13.2	12.7	9.8	17.7	10.8	6.1	5.2	11.9	15.1
Wind speed	2.70	2.93	2.98	3.04	2.86	2.51	2.65	2.77	2.49	2.50	2.55	2.54

34

734	Table 3 Statistical results of the Mann-Kendall test and Theil-Sen trend estimate for daily
735	mean values of NO_x , CO mixing ratios, temperature (T), and wind speed (WS) in
736	September and October at Dongtan (DT) site, a remote rural site near Sheshan Island
737	(SSI) site during the 2012–2017 period. The units of the calculated slopes are ppbv yr^{-1} for
738	NO _x and CO, $^{\circ}$ C yr ⁻¹ for T, and m s ⁻¹ yr ⁻¹ for WS.

	NO _x	СО	Т	WS
Slope Estimate	0.48*	2.67 [∆]	0.15 [∆]	0.21*

739 ^{*}The result is significant at the 95% confidence level.

⁷⁴⁰ ^ΔThe result cannot pass the Mann-Kendall trend test at the 90% confidence level.

741	Table 4 Statistical results of the comparisons between the simulated and observed
742	surface O_3 concentrations at Sheshan (SS), Xujiahua (XJH), Pudong (PD), DT (Dongtan)
743	and Sheshan Island (SSI) sites during September 2014. The calculated O_3 levels are
744	obtained from BC_40, BC_50 and BC_60 simulations, respectively. Values of the average
745	surface O_3 concentrations (Mean) and normalized mean bias (NMB) are displayed. The
746	NMB is defined as NMB= $\frac{\sum_{i=1}^{n} (P_i - O_i)}{\sum_{i=1}^{n} O_i}$, where P_i and O_i are predicted and observed ozone
747	mixing ratios for sample <i>i</i> , <i>n</i> is the number of total samples (numbers in parentheses).

	Cases	SS (681)	XJH (641)	PD (690)	DT (690)	SSI (720)
	Observation	39.7	30.4	40.3	46.4	57.7
Mean	BC_40	36.0	22.0	29.5	35.3	36.9
(ppbv)	BC_50	39.1	25.1	33.3	39.6	41.8
	BC_60	43.1	29.0	37.9	45.0	47.3
	BC_40	-9.4	-27.6	-26.7	-23.9	-36.1
NMB(%)	BC_50	-1.5	-17.5	-17.2	-14.5	-27.5
	BC_60	8.6	-4.6	-5.9	-3.0	-18.1

748 Figure Captions

749 **Figure 1** Land cover of Shanghai and corresponding locations and landscapes of Xujiahui

750 (XJH, urban), Dongtan (DT, rural) and Sheshan Island (SSI, remote and oceanic) stations.

751 Figure 2 Monthly wind rose diagrams averaged over the period of 2012 to 2017 at

752 Dongtan (DT) site, a remote rural site near the Sheshan Island (SSI).

753 Figure 3 Monthly mean CO mixing ratios under north (N), northeast (NE), east (E),

southeast (SE), south (S), southwest (SW), west (W), northwest (NW) and calm (C) wind

conditions at Dongtan (DT) site, a remote rural site near the Sheshan Island (SSI) during

756 **2012 to 2017**.

Figure 4 Monthly and year-round mean diurnal variations of O_3 (ppbv) at Sheshan Island

(SSI, remote and oceanic) and Xujiahui (XJH, urban) sites during 2012 to 2017.

759 Figure 5 Calculated monthly mean ratios of daily maximum O₃ concentrations (O_{3-max}) to

minimum O_3 concentrations (O_{3-min}) at Sheshan Island (SSI, remote and oceanic) and

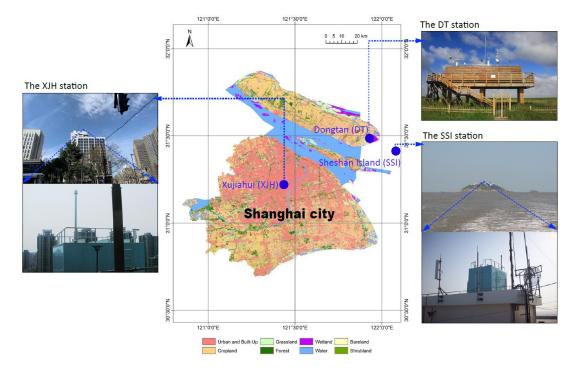
761 Xujiahui (XJH, urban) sites, respectively during 2012 to 2017.

Figure 6 Variations of (a) monthly mean O₃ concentrations at Sheshan Island (SSI, remote and oceanic) and Xujiahui (XJH, urban) sites during the period 2012–2017, (c) corresponding variations of daily mean O₃ concentrations at SSI and XJH in September and October, and (c) variations of mean O₃ concentrations during daytime (10:00-16:00 LST) and nighttime (23:00-04:00 LST) at SSI.

Figure 7 Daytime and nighttime mean O_3 mixing ratios (ppbv) at Sheshan Island (SSI) and NO_x mixing ratios (ppbv) at Dongtan (DT) site, a remote rural site near SSI under north (N), northeast (NE), east (E), southeast (SE), south (S), southwest (SW), west (W),

770	and northwest (NW) wind conditions in MAM (March-May), JJA (June-August), SON
771	(September-November), and DJF (December-February), respectively during 2012 to
772	2017.

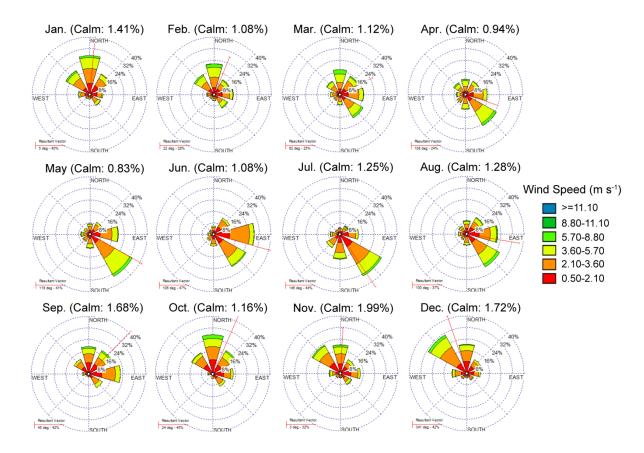
773	Figure 8 Calculated distributions of monthly mean O_3 concentrations (shades, ppbv) from
774	BC_40, BC_50 and BC_60 simulations, respectively in September 2014. Model results
775	are compared with observed mean O_3 concentrations (circles, ppbv) obtained from
776	Sheshan (SS), Xujiahua (XJH), Pudong (PD), DT (Dongtan) and Sheshan Island (SSI)
777	sites. Also shown is the calculated wind field (m s ⁻¹) averaged over the same period.
778	Figure 9 Mean differences in surface O_3 concentrations (ppbv) simulated with different
779	chemical boundaries: (a) BC_50 minus BC_40, (b) BC_60 minus BC_40, and (c) BC_60
780	minus BC_50 in September 2014. Also shown is the calculated wind field (m s ⁻¹) averaged
781	over the simulation period.



783 Figure 1 Land cover of Shanghai and corresponding locations and landscapes of Xujiahui

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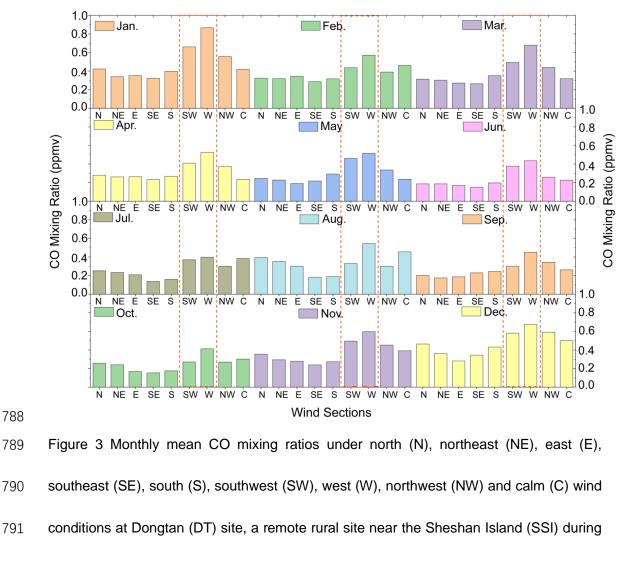
784 (XJH, urban), Dongtan (DT, rural) and Sheshan Island (SSI, remote and oceanic) stations.



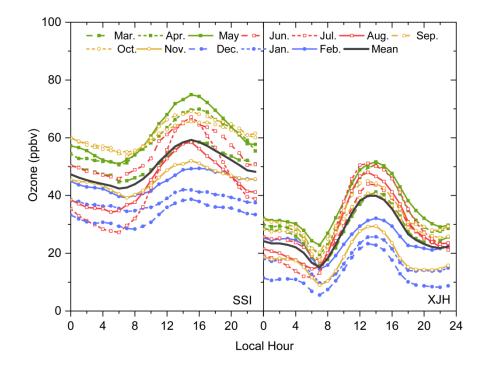
785

Figure 2 Monthly wind rose diagrams averaged over the period of 2012 to 2017 at

787 Dongtan (DT) site, a remote rural site near the Sheshan Island (SSI).



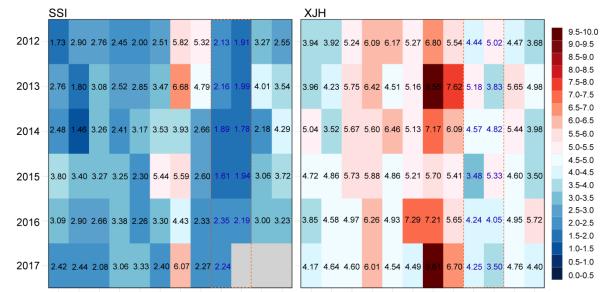
2012 to 2017.



793

Figure 4 Monthly and year-round mean diurnal variations of O₃ (ppbv) at Sheshan Island

795 (SSI, remote and oceanic) and Xujiahui (XJH, urban) sites during 2012 to 2017.



796

Jan. Feb. Mar. Apr. May Jun. Jul. Aug. Sep.Oct. Nov.Dec. Jan. Feb. Mar. Apr. May Jun. Jul. Aug. Sep.Oct. Nov.Dec.

Figure 5 Calculated monthly mean ratios of daily maximum O_3 concentrations (O_{3-max}) to

minimum O_3 concentrations (O_{3-min}) at Sheshan Island (SSI, remote and oceanic) and

799 Xujiahui (XJH, urban) sites, respectively during 2012 to 2017.

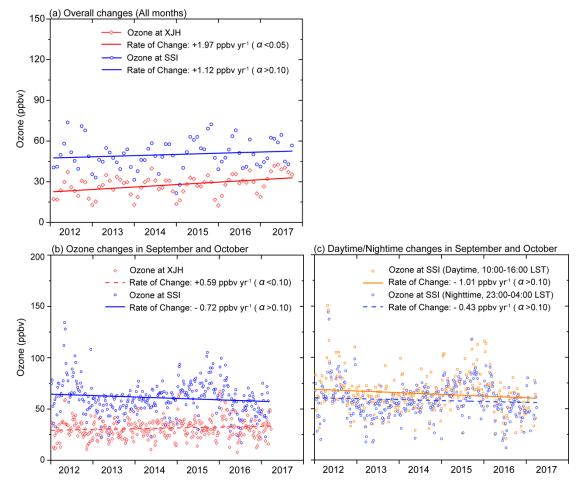
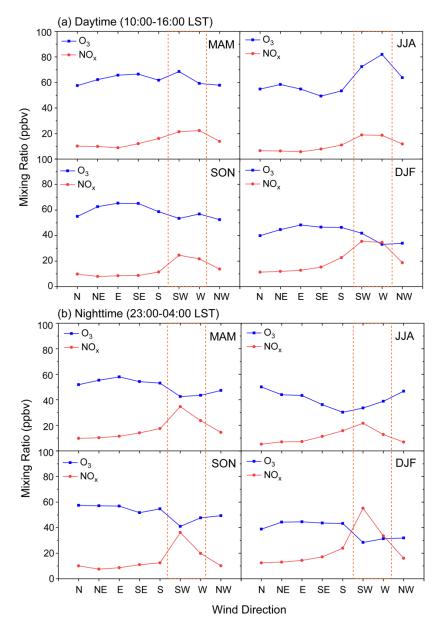


Figure 6 Variations of (a) monthly mean O_3 concentrations at Sheshan Island (SSI, remote and oceanic) and Xujiahui (XJH, urban) sites during the period 2012–2017, (b) corresponding variations of daily mean O_3 concentrations at SSI and XJH in September and October, and (c) variations of mean O_3 concentrations during daytime (10:00-16:00 LST) and nighttime (23:00-04:00 LST) at SSI.

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Figure 7 Daytime and nighttime mean O₃ mixing ratios (ppbv) at Sheshan Island (SSI) and NO_x mixing ratios (ppbv) at Dongtan (DT) site, a remote rural site near SSI under north (N), northeast (NE), east (E), southeast (SE), south (S), southwest (SW), west (W), and northwest (NW) wind conditions in MAM (March–May), JJA (June–August), SON (September–November), and DJF (December–February), respectively during 2012 to 2017.

