Reply to interactive comments on "Wildfire smoke in the lower stratosphere identified by in situ CO observations", by Joram J. D. Hooghiem et al.

September 16, 2020

General

We thank the two reviewers for their feedback on our manuscript. The manuscript has been improved by making use of their suggestions, most notably those concerning the methods section. Our replies can be found in blue, below specific comments of the two separate reviewers. In places, the revisions from the manuscript are copied, and depicted in boldface.

We should mention that we have found a mistake in one of our calculations that led to errors in several places in the manuscript. It concerns the mole fraction and isotopic composition calculations discussed in Sections 3.5 and 3.6. It does not affect the general conclusions. The ranges in 3.5 changed from $-27.5\ ‰\ \leq \delta^{13}C \leq -25.0\ ‰$ to $-27.5\ ‰\ \leq \delta^{13}C \leq -26.6\ ‰$, and $10\ ‰\ \leq \delta^{18}O \leq 16\ ‰$ to $10\ ‰\ \leq \delta^{18}O \leq 14\ ‰$. The estimated fraction of tropospheric air in 3.6 changed from 34 to 45 %.

In addition to the changes discussed below, a few additions to the manuscript have been made. A panel has been added to Figure 3 that shows the complete trajectories; this panel aids in the explanation of the methods. Also, the tropopause altitude is now added to Figure 1.

General Comments Referee #1

The paper "Wildfire smoke in the lower stratosphere identified by in situ CO observations" presents new data collected in the stratosphere by two separate sampling systems. The scientific merit is high, and the analysis methods are sound. In particular, the authors have presented an algorithm for assessing the impact of OH on CO during transport to the stratosphere. The paper overall is well structured. However, at times the sentence structure and descriptions of the methods are confusing, and several key areas in the results and discussion sections lack needed clarity/explanations. Overall, the paper is excellent, and will make a great contribution to ACP after several minor revisions (see below). Overview of Revisions Needed:

1. In section 2.2.2 and 2.2.3, the LISA sample transfer and CO isotopic measurements are described. Hooghiem et al. (2018) and Pathirana et al. (2015) are cited for reference on the LISA sampler and the CO isotopic analysis line, respectively. However, I was unable to find tests in either reference that show storage tests performed on the sampling containers used in this study. The authors make special note to point out the storage flasks are sealed with Viton O-rings, which is known to contaminate for CO mole fraction on the order of 1-2 ppb per day (e.g. Novelli, Steele, and Tans, 1992). A reference or supplemental section should be included that describes laboratory tests which demonstrate that $x(CO)$, $\delta^{13}CO$, and $\delta^{18}CO$, are unaffected by a 7-month storage time in flasks with a Viton oring. This is especially important given the low mole fractions you have observed in the stratosphere as well as the few data points

you present here. Any sampling bias could have a large impact on your source attribution.

A stability test was performed using stratospheric air samples obtained from the LISA flights over Sodankylä in 2018. Analysis, transfer, and storage were performed using the same methods as presented as in the present study. Re-analysis on 4 samples was performed using the same Picarro instrument in July 2020. The drift in $x(\text{CO})$ was computed to be 0.05 ppb day^{-1}, assuming a linear drift.

The stability of the $\delta^{13}\text{C}(\text{CO})$ and $\delta^{18}\text{O}(\text{CO})$ in these flasks was not assessed directly. The flasks used for storage have a two O-ring configuration which is known to give better results in storage test (Sturm et al., 2004), this is added to the methods section: **"The flasks have a Rotulex connection and are sealed with two Viton-70 O-rings, providing better sample stability than single O-ring configuration (Sturm et al., 2004)."** As discussed in text, the stratospheric background measurement compared well to other stratoshperic measurements, and the plume sample differs significantly.

A computation is provided to show the likely magnitude of a contamination. The source of contamination shall be the air surrounding the flasks. Assuming that the contamination can be modelled using the mixing model provided in the manuscript, e.g. Eq. 1 and 2. Atmospheric measurements on northern hemisphere background air from Mak et al., 2003 provides values $x(\text{CO}) = 100$ ppb, $\delta^{18}\text{O} = 0$ ‰, and $\delta^{13}\text{C} = -29$. The estimate results in no observable difference in $\delta^{13}\text{C}$, and -0.7 ‰difference for the plume in $\delta^{18}\text{O}$. Alternatively, more polluted, air values like one would find in spring $x(\text{CO}) = 150$ ppb, $\delta^{18}\text{O} = 4$ ‰, and $\delta^{13}\text{C} = -26$. Also leads to less than 0.2 ‰for both $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values of both the plume and background.

In this computation, fractionation is ignored entirely. Also, the computation is very sensitive to the contamination source values and drift rate. Here we tried to give a plausible estimate based on atmospheric values. It should be noted that a drift in mole fraction does not necessarily mean a change in isotopic composition (as illustrated above) nor does a stable mole fraction guarantee a stable isotopic composition during storage. A systematic bias cannot be excluded, but may be small. We have explicitly stated in the method section that we cannot exclude a bias due to the long term storage:

"Yet, it can not be excluded that the isotopic measurements, see below, are biased by more than a ‰."

2. The analysis of the enhancement CO:CO$_2$ ratios presented in the results and discussion is confusing as it is presented currently. In the methods section, the authors describe a common method for obtaining enhancement ratios: assuming that an air parcel is a mixture of some sources plus a background. The authors further show that they have made good background measurements for the altitudes of interest in the stratosphere on Sept 6 and 7, at least for mole fraction. However, in the analysis, particularly in Figure 4, it does not seem that any background values have been subtracted off. The enhancements

presented in figure 4 appear to just be the mole fraction measurements from 4 and 5 September in figure 1. In figure 4, the authors appear to be attempting to find a slope to predict the overall $CO:CO_2$ ratio in the plume. The x axis should range from 0 to 1 (or possibly 1.5) ppm, consistent with the 6th and 7th September xCO_2 curves (background) being subtracted from the 4th and 5th September xCO_2 curves. The xCO enhancement on the y-axis should similarly be adjusted. Given that the background curves are not constant through the altitude range considered, I think that the subtraction may alter the results of the slopes presented in figure 4. This is, of course, provided that the error is not simply a typographical error in the labeling of the axis. If, on the other hand, the authors are attempting only to show the $CO:CO_2$ ratios of the plume, and not subtract the background, then the methods should be clarified.

A description of the methodology applied was missing. Two paragraphs have been added to the methods section to clarify the methodology:

"CO and CO_2 are co-produced in burning processes, and their emissions into the atmosphere result in an enhancements, ΔCO and ΔCO_2, compared to background air. The enhancement ratio of ΔCO : ΔCO_2 is typically high for wildfires, and decreases over time due to photochemical loss of CO (Mauzerall et al., 1998). The enhancement ratio is conserved during mixing, if the background is constant. Thus, in this case, the enhancement ratio can be directly obtained as the slope of a linear regression performed on AirCore CO and CO_2 data for the two separate AirCore flights that sampled the plume on 4 and 5 September 2017."

"Alternatively, measurements of background air can be used to compute the enhancement ratio directly. It is assumed that the flight performed on 6 September is representative for the background air adjacent to the plume. CO and CO_2 data are smoothed using a moving average with an averaging window of 25 data points in order to reduce the analyser noise in CO. The background air is interpolated on isentropic surfaces to the observed plume altitude, and the enhancements are computed directly on isentropic surfaces. As we will see, the total enhancement in CO_2 is small, and only the results with $x(CO_2) > 0.2$ ppm are used in the subsequent computation of the enhancement ratios."

The enhancement ratio is preserved if a background has constant mole fractions in the tracers for which the enhancement ratio is calculated. So the implicit assumption made in the manuscript, was that the background is constant. If the background is constant, it does not matter which dimension is used to intersect the plume (i.e. horizontal or vertical). Moreover, the actual background mole fractions need not to be known in order to determine the enhancement ratio. Subtracting a constant value for both $x(CO_2)$ and $x(CO)$ respectively would merely result in a change of the intercept of the linear fit applied. The intercept is not of interest. Hence, when the assumption of a constant background can be justified, the enhancement ratio of the plume is the same as the slope of a tracer-tracer correlation. The justification is provided by the overall regression

statistics, the $r^2 > 0.83$ and the 95 % confidence interval on the slope.

We agree that validity of the assumption of constant background is questionable. This is now discussed together with an additional computation of the enhancement ratios (see added paragraph above). The discussion is updated with the following paragraph

"The enhancement ratios computed directly are much higher than the results discussed above. First, it is difficult to assume what background should be used. The stratospheric background mole fractions, especially the CO_2, varies with altitude and in time, e.g. comparing the AirCore profiles from 6 and 7 September, which questions the assumption of a constant background. It is thus difficult to state with certainty that data from 6 September is representative for the background. However, as the plume, and air directly adjacent to it, move together, it can be assumed that the air surrounding the plume has constant mole fractions. The best estimate of the mole fractions are those measured above and below the plume, obtained from the vertical profile. Hence, the most reliable estimate of the enhancement ratio is obtained from the regression analysis discussed above. On a final note, the very small standard deviation of 1–2 ppb obtained for 4 September is largely due to the small amount of data, $N = 5$, and the large averaging window of 25 data-points."

3. Lastly, the LISA flight from the 4th of September is mentioned in the methods section, but I have been unable to find this data in the results and discussion of the isotopic measurements. Why is this missing? It would further strengthen your source attribution section as you currently only have one plume and one background point. If it was discarded, the reasoning should be presented. I recommend that you add this point into your dataset if possible.

It is correct that no data is presented here from the LISA flight of September 4. In the methods a more generic description of data retrieval is presented. Data collected from LISA on 4 September has no further relevance to this study, since it didn't collect samples in the lowermost stratosphere. The date of 4 September is mainly relevant for the AirCore measurements of which contains the plume measurement. A line has been added into the methods that clarifies which data from LISA is used in the presented study.

"The AirCore was analysed for CO_2, CH_4, and CO mole fractions, for details see Section 2.2.1. In this work only the AirCore CO_2 and CO profiles are used. LISA samples have been analysed for CO_2, CH_4, and CO mole fractions, see Section 2.2.1, and the stable isotopic composition of CO, see Section 2.2.3. Here only the LISA CO mole fraction and CO isotope measurements at the plume altitude, see Section 3.1, are used, one sample from 5 and 1 sample from 6 September 2017. Although measured, LISA CO_2 appears to suffer from a bias as concluded from comparison with AirCore measurements, see Hooghiem et al., 2018."

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Specific Comments Referee #1

Pg 1, Ln 16: 1km? Into the stratosphere or you only sampled 1 km of the plume? Clarify.

The plumes vertical extent is 1 km, this has been clarified in the revised version.

"Finally, the plume was extending over 1 km in altitude, as inferred from the observations."

Pg 2 Ln 24: Specify dates.

The dates of the lidar observations have been added.

"In 2017, a large smoke plume in the stratosphere was observed on several days between 24 August and 26 September by ground-based LIDAR and the Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) aboard the CALIPSO satellite (Khaykin et al., 2018)."

Pg 2 ln 47: You say that Mauzerall et al. 1998 also measured $CO:CO_2$ just before stating that only Jost et al measured CO_2 allowing for $CO:CO_2$. Please reword/clarify this statement.

It was meant to say, that Jost et al presented the only observation of the enhancement ratio in the stratosphere. This is clarified.

"Of the stratospheric observations, only Jost et al., 2004 measured CO_2, allowing $\Delta CO : \Delta CO_2$ to be quantified, confirming the smoke's origin."

Pg 3 ln 69-71: These three sentences need to be reworded. They are awkward to read. Explain what you mean by the sample is limited. I.E. explain how the low pressure and high rate of descent in the stratosphere results in 0.3L air for xxx m of stratosphere, compared to 1.1L of air for xxx m of troposphere. The low resolution and small sample volume lead to the need for LISA, etc...

The sentence have been revised for a better explanation.

"The AirCore has a volume of 1.4 l. The AirCore takes a sample passively, and due to the low pressure in the stratosphere only about 0.3 l of the sample is stratospheric. The AirCore has a vertical resolution of 374 m at 200 hPa or 12 km altitude."

Pg 3 ln 73: Larger amount relative to AirCore? You specify the L STP for AirCore, you should state the same here for LISA.

We have added the total amount of stratospheric air that LISA samples for a direct comparison.

"The active sampling results in a larger amount of sample 180–800 ml per sample, thus allowing for isotope analysis."

Pg 3 ln 75: Compare to AirCore here as well.

We have added the vertical resolution of 374 m at 12 km altitude of the AirCore used in this study (see reply to the comment about "Pg 3 ln 69-71" above.

Pg 3 ln 80: "In addition to . . . " Combine these two sentences

The last 2 sentences have been combined. The contribution of the radiosonde is now summarized in one sentence.

Pg 3 ln 87: Explain why two different instruments and models for the two analyses?

Two instruments were used in order to reduce time between payload retrieval and analyses for both AirCore and LISA in the field. The model only matters in the sense that they both measure CO_2, CH_4, CO, and H_2O. We have added this to the revised manuscript.

"Two analysers were used, to allow for simultaneous analyses of both AirCore and LISA samples after payload retrieval (analyser models used: Picarro G2401 (AirCore), Picarro G2401-m (LISA) see e.g. (Crosson, 2008) for more information on the CRDS-analyser)."

Pg 3 ln 84 and throughout text: Note: mole fractions to be represented by lower case x with subscript chemical formula per IUPAC standards. I personally do not have a problem with the nomenclature used in the text, but the authors may wish to make the change for technical correctness

The manuscript, both revised and original, were intended to be fully compatible with definitions by IUPAC (Cohen et al., 2007) at least as far as equations go. For example Equation 1, correctly uses italic x for the mole fraction with upright text to specify its origin. In text we believe "CO_2 mole fractions" as well as its symbolic representation "$x(CO_2)$" are allowed, just as you would say pressure measurements in air vs p measurements in air. Note that footnote two, chapter 2.10 on page 47 of Cohen et al., 2007, states: ". . . When the chemical composition is written out, parentheses should be used, $n(O_2)$", which also applies to mole fractions, according to page 48 of Cohen et al., 2007.

Pg 4: Section 2.2.3: I am not seeing the number of samples you have measured. You state above that you flew LISA on 3 flights, with a total of 4 samples per flight. However, you only present 1 data point per day. Please explicitly state the number of flasks you analysed for stable isotopes and if you averaged them for a given day. This is especially important below when you do not show the 4th Sept data (See further comments below and above on 4-Sept data).

This has been clarified. See also our reply to main comment number 3.

Pg 5 ln 124: Replace "concentrations" with "mole fractions".

We have replaced "concentrations" with "mole fractions".

Pg 5 Ln 135-138: It is unclear here what aspects of the CLaMS you used? This sentence seems to indicate that you used the advection scheme but not the mixing scheme, while at the same time implying that both the advection and mixing schemes are needed to resolve 2d filamentary structures. Please clarify.

We only used the trajectory module of CLaMS. A complete reconstruction of the 2-d (or even 3-d) structures of the exhaust would be much more complicated. Also, little would be known about the vertical velocity. Therefore we added the comparison with the CALIOP data from which we could piecewise trace the observed cloud. We have separated generic information from what this work uses

"In the present study, only the advection scheme is used. First, information on the aforementioned additional vertical motion is lacking. Secondly, only synoptic scale transport is of interest to determine the source region of the smoke."

Pg 5, ln 147: How accurately? This paragraph does not indicate the scale to which the smoke plume can be traced using this method. Suggest removing this word.

This word has been removed.

Pg 6 ln 153: "similar balance" should be "similar mass balance".

Changes implemented as suggested by the referee.

Pg 6 ln 155: Suggest modifying equation (2) to include "approx. equal" since delta notation here is a good approximation for small changes in delta, but strictly speaking, there is loss of tracer (very small) in this approximation (see Tans, 1980).

The $\approx$ is adopted in the revised manuscript, and Tans, 1980 is added as a reference.

Pg 6, ln 170: Remove the sentence "The production ... neglected". This is confusing here and is not relevant to this section which is explaining the oxidation of CO by OH. This can be explained further in the Discussion section about how you determined the actual source.

The sentence is removed, as suggested.

Pg 8 Ln 200: "Volumetric Air Fraction": this is a confusing term. Looking at your equations, f is the fraction of the total mole fraction. You have a measured mole fraction for your given air parcel that you are trying to partition with your model. Furthermore, you are looking at CO here, not air, and volumetric is

additionally confusing as you have introduced no volumes in this equation that I can determine. Please clarify or restate.

f is the fraction of molecules coming from different sources substantiated by the subscripts (tropospheric, stratospheric, or wildfire). We have added a clarification for the parameter f defined in text. From the ideal gas law it follows that at constant temperature and pressure f can be computed equivalently from volume or from number density. Yet, we agree that the use of the term "volumetric" might be more confusing. Therefore we have redefined it to fraction of molecules.

"Here $f_{\mathbf{bg}}$ is the fraction of molecules of bg in ap, and similar for src"

Pg 8- Pg 9, Ln 216-219: This needs more explanation. Please expand or clarify. This currently leaves the reader with the impression that your signal to noise ratio is something like 1:1, which would mean that you could not differentiate between the measurement noise of your measurement system and real atmospheric signals. This is certainly not the case for the AirCore. Also, please explain/expand on why 2x the measurement uncertainty was chosen and why this is assumed to be valid. The current level of explanation makes this choice seem arbitrary.

Pg 10 Ln 245-247: As with my previous comment on this section, you need to clarify and explain more clearly how you are assigning uncertainty, how you are interpreting your results within the bounds of that uncertainty. Why are you limiting your results to live within the 1 sigma bounds, for example? I'm not suggesting this approach is wrong, but it certainly needs further explanation.

This reply concerns the two previous comments together. The methodology has been updated with a better clarification and argumentation, as where the inputs of the Monte-Carlo simulation come from, including the chosen uncertainties.

"The stratospheric end-member definition and the plume observation are based on the balloon-borne observations by LISA, outside and inside the plume, respectively, see Table 4. The plume mole fractions are assigned an uncertainty equal to the measurement uncertainty of LISA. The uncertainty of the stratospheric end-member is also a good measure of the variability surrounding the plume based on AirCore measurements between 12 and 14.5 km of 6 and 7 September. The uncertainty of the isotopic composition in the Monte-Carlo simulation is set equal to twice the measurement uncertainty (see Section 2.2.3) because the stratospheric variability of the stable isotopic composition of CO is unknown. Secondly, there may have been a small drift in the isotopic composition."

Figure 1: Why is 4-Sept LISA measurement missing from this plot?

It is not missing, on that day there was no sample in the shown altitude

range. Note that with the updated methodology, it is made clear that LISA data from 4 September is not used throughout this work.

Pg 13 Ln 276-279/throughout section 3.3/Figure 4: Figure 4 and the description presented here are confusing (see General comment 2 above). You refer to enhancements and enhancement ratios, and in Section 2 you describe the source attribution in terms of sources on top of a background, or wildfire plume injection on top of the background stratosphere. However, in this section, it is not clear where you have subtracted your background or what you are using for a background. Are you using the profiles from the 6th and 7th averaged to assume a "typical" stratospheric background CO and CO_2? Furthermore, figure 4 would be clearer if you displayed the enhancement CO vs the enhancement CO_2 (e.g. the background subtracted off each parameter).

Here we like to refer to our reply to the main comment.

Pg 15 Ln 308: Given your stated measurement uncertainty, can you really say this is a significant change?

Since the difference of 0.8 per mill is larger than the measurement uncertainty mentioned in the methods section (0.5), it is statistically significant. Yet, Ln 308 tried to state that it is not sure if this is also larger than the natural variability. This is now restated.

"Though the observed difference in $\delta^{13}C$ between the plume and the background sample is significant (0.8 ‰), but could also be the result of natural variability."

Page 15 Table 4: Include data from 4-Sept if possible

As it is not used, see earlier replies concerning LISA measurements from 4 September; it is not added.

Pg 19, ln 376: Again, you are not presenting the enhancement ratios here. This is just the $CO:CO_2$ ratio from your overall flight. To get the enhancement ratio, you would subtract your background profile from your plume profile to plot the enhancement $CO:CO_2$ ratios.

Here we also like to refer back to the replay to the main comment.

General comments Referee #2

The paper "Wildfire smoke in the lower stratosphere identified by in situ CO observations" presents measurements of CO mole fractions, CO_2 mole fractions, and isotopic composition of CO from two balloon-borne instruments. Overall, the paper presents new datasets and includes in depth analysis of the measurements. It is generally well written, with clear discussion of results. The

paper would make a good contribution to ACP, after the following comments are addressed.

For the most-part, I found the description and justification for the methods chosen were complete and well-referenced. However, I do not have the expertise to fully evaluate some details of the data collection and analysis methods (Sect. 2.1, 2.2 2.3.2, 2.3.3). In some sections, I found it a bit difficult to follow the methods. For example, some descriptions of the methodology are written into figure and table captions. Throughout the paper, it would be helpful if high level information was provided at the beginning of each section to help guide the reader and explain how and why certain methods were applied. I have included some specific comments below, which include examples where additional descriptions would help.

We thank Anonymous Referee 2 for the kind words of appreciation for our manuscript. We hope to have addressed the issue of readability of the methods section by incorporating the essential information into the main text. Specifically, information that was only included in the captions of figures and tables.

Specific comments from Referee #2

Line 15: "The in situ observations provide information ... of the 2017 smoke plume" The closing sentence of the abstract is a bit unclear. Is the "information on the trace gas chemistry" what was described in the previous part of the abstract? If so, perhaps could use more specific language (e.g., what new information is provided by this study?). Also, I don't see much mention of the 1 km width of the plume in the text (I assume this inferred from Sect. 3.1?). If this is a key result, then maybe it warrants more discussion in text.

We have changed the closing statement of the abstract, to remove ambiguity. We have added explicitly to section 3.1 that the vertical extent of the observed plume is 1 km. Something that in the earlier version was left for the reader to determine from line 250 of the initial manuscript, as pointed out by the referee. We have repeated this observation in the conclusions.

Revised in the abstract:

"Finally, the plume was extending over 1 km in altitude, as inferred from the observations."

Added in the results section:

"The plume, extending over roughly 1 km in altitude, was well above the tropopause . . . "

Added in the conclusions:

"present on two consecutive days, and extending over 1 km in altitude."

Line 59 "Methods": A very short overview of what the AirCore vs LISA measure (mole fractions vs isotopic composition) and to what approximate vertical / temporal resolution would be helpful here.

This information has been included in the revised manuscript. See also the reply to the main comment 3 from referee 1 above.

Line 139: I found it very difficult to understand the method used for the back trajectories until I reached Fig 2 and the associated text in Sect. 3.2, which walks the reader through this process. Until reaching this figure, I found terms like "piece-wise manner" confusing. To remedy this, the authors could merge Sect. 2.3.1 directly into Sect. 3.2 or perhaps make Sect. 2.3.1 a bit more general with a forward reference to Sect. 3.2 (e.g., something like used back trajectories from CLaMS and a piecewise method illustrated in Sect. 3.2).

The second suggestion is adopted in the revised manuscript. Also, the text in Sect. 2.3.1 has been changed, to clarify the method used.

"Therefore, a correction for the vertical displacement is determined in correspondence with CALIOP elastic backscatter-ratios at 532 nm (Winker et al., 2010) as illustrated in Section 3.2."

Line 148: A bit of an overview would be helpful for Sect. 2.3.2 and 2.3.3. How are the methods being applied to the data? What information do the methods in Sect. 2.3.2 provide compared to Sect. 2.3.3? Have these methods been applied to similar datasets before?

We have added a statement to both Sect. 2.3.2 and 2.3.3 to clarify what information is retrieved from the data. We have added a few references where the reader can find examples of earlier application of the method specified Sect. 2.3.2 (added references: (Vimont, Turnbull, Petrenko, Place, Sweeney, et al., 2019; Vimont, Turnbull, Petrenko, Place, Karion, et al., 2017; Gromov and C. A. Brenninkmeijer, 2015)). The method in Sect. 2.3.3 is, to our best knowledge, not used in atmospheric science, and we therefore cannot provide a good example. The method in Sect. 2.3.3 is essentially an extension to the two end member case presented in Sect. 2.3.2. Sect. 2.3.3 has been revised, to improve clarity.

To section 2.3.3 (used to be 2.3.2) this sentence is added

"The method usually employed to determine the source signature of an observed pollution, is to assume that a measured air parcel is the result of mixing of background air and a polluting source, i.e. two end-member mixing."

To section 2.3.4 (used to be 2.3.4) this sentence is added

"The mass balances in Equations 1 and 2 can be extended to allow

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for more than two end members"

Line 152: Did not define "f" in the text. Check throughout that all variables in equations are defined. In the revised manuscript f is now defined. We have also taken a good look at other definitions of variables.

Line 191: I found this section confusing at first. I had to jump around the text to understand how Table 3 was constructed and how this information was used. Some of the methods/categories are only really described in the Table 3 and Figure 6 captions. Also, the Table 3 caption references "Monte-Carlo simulation" at the very top of the caption, but the MC simulation is not mentioned until much further in the text. One possible way to improve this would be to give a high level overview when the table is first introduced. This would describe why the table was put together, how the various lines of the table were compiled (with cross-references to relevant sections), the purpose / input parameters of the MC simulation when the table is first introduced.

Table 3 caption: The caption is a bit confusing –review/rewrite and maybe move some of the details into the text describing the methods. For example, "...does not significantly affect the results" – the results of what? The MC simulation?

Table 3: Where did the numbers for AP OH-Corrected come from? Are these related to the numbers given around line 337 – if so, how exactly were they chosen (they don't seem to exactly match any of the numbers given in the text)?

A paragraph is added to clarify the purpose of Table 3 and the origin of the variables displayed in Table 3. The authors agree, upon rereading the caption of Table 3 that it was unclear, and thus the caption has been revised. This information has also been put in the main text. Given the size of the revision, the authors refer to the revised manuscript, Section 2.3.4.

Line 237: "This thus differs . . . in Table 2". Could you clarify this statement? It looks like the values in Table 3 fall within the range of Table 2 – but are a narrower part of the range (because from specific type of biomass burning?)

The statement has been changed. It is indeed a narrower part of the range specified in Table 2.

"This is thus a subset from the range of signatures displayed in Table 1, which provide a more general summary"

Line 240: Please add a bit of a description of the Monte-Carlo simulation. It's not clear what data is being used, what equation it's applied to, and what the desired output is until the reader gets to the caption of Fig 6.

We have added a paragraph to include this information in the methods section, see the revised manuscript 2.3.4.

Line 255: "The observed CO_2 ... Which allowed for determination of enhancement ratio" Is the observed CO_2 difference significant for both LISA and AC? It looks like LISA is biased high compared with AC (e.g., non-enhanced day for LISA is comparable to biggest enhancement for AC). Also, why does the slight increase in CO_2 allow for enhancement ratio calculations?

Unfortunately, LISA samples appear to suffer from a bias in CO_2, see the LISA and AirCore comparisons in (Hooghiem et al., 2018). This is now stated in text Sect. 2.2.1. LISA CO_2 data was not used in the analysis presented by this manuscript. It was initially displayed for completeness of the data description. On a second thought, CH_4 was also measured by both LISA and AirCore. Yet this data has not been used nor presented in this manuscript. Therefore the CO_2 measurements of LISA have been removed from Fig. 1b and removed from Table 4, as they don't play a role in further analysis.

Line 259: This section could use an opening sentence describing what the back trajectories are being used for.

This section has been updated according to earlier comments.

Line 260: What is the "match distance"? Is this the minimum distance from the back trajectory to the CALIOP scan? Also, since the centre/upper altitude are not shown – is this the lower altitude? Is this starting from 13.3 km (based on Sect. 2.3)?

The Match distance is the horizontal distance of the trajectory location to the orbit track at the time of the satellite overpass. The distance of 250 km was chosen to balance between a coverage of the plume and a sufficient number of matches. It should be well below the geographical extent of the wildfire plume in order to follow the plume. The trajectory shown, was the trajectory with the lowest match distance. This was the lower trajectory of the 4 September observation, which is stated in the manuscript.

"Wherever the distance was smaller than 250 km"

And to the caption of Figure 2:

"A match on 3 September between CALIOP and the lower CLaMS back-trajectory starting on 4 September"

Figure 2 caption: "...the correspondence was sufficiently good". Does this mean the altitude correspondence between the back trajectory and CALIOP aerosol enhancement?

It means that the computed back-trajectory location was coinciding with aerosol enhancements. We have changed the text to be more exact. In the manuscript only locations where the CLaMS trajectory location was not at an observed aerosol cloud are shown. A line has been added into the main text to clarify this.

"Other matches between back-trajectory location and CALIOP overpass was always at a layer of aerosol enhancement and no new back-trajectories were initialized. These match results are not shown."

Line 277: How are the slope / uncertainty calculated?

The slope was computed using linear least-squares regression. The uncertainty is the standard error for the gradient. For background, the routine used for the computation is provided by python scipy package (scipy.stats.linregress). We have updated the manuscript using the more common 95 % confidence interval for the slope, which can be computed from the standard error of the gradient.

Line 307: How do you know that the plume is clearly not stratospheric? Or is it just clear that the plume is different from background?

The text has been revised, and now states that the plume is an enhancement compared to the background and has unusual high mole fractions of carbon monoxide:

"clearly shows that the CO plume is different than the background"

Line 332: How did you choose the ranges of possible values from Fig 5?

The text clarifies where these values came from. They were defined as the range provided by the orange curve in figure 5, i.e. the computation using the value of OH deemed the most representative for the plume.

Figure 5: The numbers in the legend aren't showing up correctly. (No value for the blue marker, reading at 1^6 instead of 10^6 for the orange and green markers)

A pdf was uploaded to ACPD with correct Figures, and we have not checked the document thoroughly enough after conversion. The correct figure can be found in the marked up version as well as the revised version.

Line 343: Why is it within the range for "biomass burning" in Table 2 but not within the range for "wildfire smoke" in Table 3?

Table 3 provides a narrower range, and is a subset of the range in Table 2. Table 2 provides a summary of our knowledge of isotopic source signatures. The values in Table 3 are a subset of Table 2 because of two criteria: 1) smoke from the fire comes from a boreal forest containing mainly C3 plants 2) the fact that the smoke rose to a very high altitude (above the tropopause), shows that the combustion temperature was very high, resulting in higher δ^{18}O values; see the discussion by Kato et al., 1999. This short discussion has been added to our methods section. What is excluded in this analysis is mixing in the troposphere. The OH chemistry is much less important in the 5 hour up-draft. Qualitative this would shift the isotopes more to range in Table 3. Mixing wildfire smoke

with depleted tropospheric air leads to lower δ^{18}O and δ^{13}C values; undoing the mixing would increase them. Revised paragraph:

"In addition, the wildfire smoke signature, see Table 1, is subjected to a large variability due to the type of burned plants (categorised as C3 or C4, with a difference in photosynthesis), the burning temperature, and possibly the groundwater isotopic composition (Kato et al., 1999). Since the wildfire originated in Canada, the fuel consisted mainly of C3 plants which are typically more depleted in δ^{13}C. Furthermore, the fire was energetic enough to trigger a pyro-Cb event, which makes it reasonable to assume that it was in an efficient burning regime, which typically leads to higher δ^{18}O. Therefore, the range of isotopic composition of CO of the wildfire smoke assumed here is according to atmospheric measurements around forest fires (C. Brenninkmeijer et al., 1999), see Table 3. This is thus a subset from the range of signatures displayed in Table 1, which provide a more general summary."

Line 414: You mention that "little is known of the CO isotopic composition". Do your measurements of the background air mass contribute to this knowledge? If so may be worth mentioning the background measurements in the abstract.

We have added the measurement to the abstract as per the suggestion.

"In addition to CO mole fractions, CO$_2$ mole fractions as well as isotopic composition of CO (δ^{13}C and δ^{18}O) have been measured in air samples, from both the wildfire plume and background, . . . "

Line 426: "Yet another event was modelled . . . " is this another wildfire event? If so, specify.

Yes, indeed another wildfire event was modelled. We have updated the text with more details about this event.

"Cammas et al., 2009 modelled their observations of wildfire smoke from several fires in Canada and Alaska and they estimated the amount of polluted boundary layer air above the tropopause to be 15–20 %."

Technical corrections from Referee #2

All technical corrections and suggestions below have been incorporated into the revised manuscript.

Line 10: "Back-trajectory analysis, performed with . . . Date of 12 August 2017" Very long sentence. Consider breaking it up.

"Back-trajectory analysis was performed with the Chemical Lagrangian Model of the Stratosphere (CLaMS), tracing the smoke's origin to wildfires in British Columbia with an injection date of 12 August 2017. The trajectories are corrected for vertical displacement due to heating of the wildfire aerosols, by observations made by the Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) instrument."

Line 13: "Colombia" $\longrightarrow$ "Columbia" (also, line 426, 434)

Line 145: "Wherever the distance was below", replace with "Wherever the distance was smaller than"

Line 245: "its" $\longrightarrow$ "it is"

Figure 2 caption: Check capitalization after periods. Also, last sentence is run on "CALIOP, the time" $\longrightarrow$ "CALIOP. The time"

Line 307: ", see Table 4" $\longrightarrow$ "in Table 4"

Line 308: "stratospheric" $\longrightarrow$ "the stratosphere"

Line 309: "it cannot be excluded" $\longrightarrow$ "it cannot be ruled out"

Line 356: delete repetition? "... was used to determine fractions of tropospheric and stratospheric air in the plume"

Line 414: "pollution air" $\longrightarrow$ "polluted air"

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Wildfire smoke in the lower stratosphere identified by in situ CO observations

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Abstract. Wildfires emit large quantities of aerosols and trace gases, which occasionally reach the lower stratosphere. In August 2017, several pyro-cumulonimbus events injected a large amount of smoke into the stratosphere, observed by lidar and satellites. Satellite observations are in general the main method of detecting these events since in situ aircraft- or balloon-based measurements of atmospheric composition at higher altitudes are not made ~~frequent~~ frequently enough. This work presents accidental balloon-borne trace gas observations of wildfire smoke in the lower stratosphere, identified by enhanced CO mole fractions at approximately 13.6 km. In addition to CO mole fractions, CO_2 mole fractions as well as isotopic composition of CO ($\delta^{13}C$ and $\delta^{18}O$) have been measured in air samples, from both the wildfire plume and background, collected using an AirCore and a LIghtweight Stratospheric Air sampler (LISA) flown on a weather balloon from Sodankylä (4–7 September 2017, $67.37°$ N, $26.63°$ E, 179 m.a.s.l), Finland. The greenhouse gas enhancement ratio ($\Delta CO : \Delta CO_2$) and the isotopic signature based on $\delta^{13}C(CO)$ and $\delta^{18}O(CO)$ independently identify wildfire emissions as the source of the stratospheric CO enhancement. Back-trajectory analysis ~~,~~ was performed with the Chemical Lagrangian Model of the Stratosphere (CLaMS), tracing the smoke's origin to wildfires in British Columbia with an injection date of 12 August 2017. The trajectories are corrected for vertical displacement ~~,~~ due to heating of the wildfire aerosols, by observations made by the Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) instrument~~, trace the smoke's origin to wildfires in British Colombia with an injection date of 12 August 2017.~~. Knowledge of the age of the smoke allowed for a correction of the enhancement ratio, $\Delta CO : \Delta CO_2$, for the chemical removal of CO by OH. The stable isotope observations were used to estimate the amount of tropospheric air in the plume at the time of observation to be about ~~34 ± 14~~45 ± 21 %. ~~The in situ observations provide information on the trace gas chemistry of smoke plumes that reach the stratosphere, as well as the vertical extent of~~ Finally, the plume was extending over 1 km ~~of the 2017 smoke plume. add a line to the abstract~~ in altitude, as inferred from the observations.

1 Introduction

Wildfires emit a large quantity of polluting trace gases and aerosols into the atmosphere (Crutzen and Andreae, 1990; Andreae, 2019). These trace gases and aerosols affect the radiative transfer properties of the atmosphere and lead to the formation of tropospheric ozone. Not only the troposphere is affected, but the smoke also occasionally reaches the lower stratosphere (Waibel et al., 1999; Fromm et al., 2000; Fromm and Servranckx, 2003; Jost et al., 2004; Fromm et al., 2010), enhancing aerosol levels and ozone (Fromm et al., 2005), with potential global effects (Peterson et al., 2018).

In 2017, a large smoke plume in the stratosphere was observed ~~on several days between 24 August and 26 September~~ by ground-based LIDAR and the Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) aboard the CALIPSO satellite (Khaykin et al., 2018). This smoke was attributed to Canadian forest fires, injected by pyro-cumulonimbus (pyro-Cb) events. The cumulative smoke mass injected into the stratosphere by five distinct pyro-Cb events was estimated to be 0.1 to 0.3~~Tg~~ Tg (Peterson et al., 2018). The smoke mass density was further characterized using the Aerosol Robotic Network (AERONET) and Moderate Resolution Imaging Spectroradiometer (MODIS) (Ansmann et al., 2018), and the micro physical properties of the smoke were determined by LIDAR studies (Haarig et al., 2018; Hu et al., 2019; Baars et al., 2019).

Past injections of wildfire smoke into the stratosphere were mainly identified and characterized using satellite observations (e.g. Fromm et al., 2010). Nevertheless, wildfire smoke has been observed from in situ aircraft measurements as well. First, Waibel et al. (1999) reported a CO-plume in the Northern Hemisphere (NH) extra-tropical lowermost stratosphere at 10 km altitude. The plume was associated with the extensive 1994 burning season. In addition, Hudson et al. (2004), Ray et al. (2004), and Jost et al. (2004) found several smoke layers between 14.7–15.8 km ($\theta = 368$ to 393 K). The enhanced levels of CO, up to 193 ppb, were found in the NH subtropical lower stratosphere (25° N), which was 1.3 km above the local tropopause. They attributed the origin of the smoke to North American forest fires. Finally, Cammas et al. (2009) reported on the injection of a smoke plume into the stratosphere also associated with North American forest fires.

In situ observations of wildfire smoke are typically identified by an increase in mole fractions of CO (Waibel et al., 1999; Jost et al., 2004; Cammas et al., 2009). In addition to CO, Cammas et al. (2009) measured O_3, NO_x, and PAN. These measurements correlate well with CO and are thus additional tracers for wildfire smoke. Furthermore, Hudson et al. (2004) and Jost et al. (2004) measured particle mass spectra containing carbon, potassium, organics, and ammonium ions. The stratospheric particle mass spectra were compared to mass spectra obtained from direct smoke measurements in the troposphere (Hudson et al., 2004), confirming the presence of smoke in the stratosphere.

Wildfire smoke has distinct trace gas source signatures. One way to identify the source of smoke is by using the enhancement ratio of $\Delta CO : \Delta CO_2$ (Mauzerall et al., 1998), and another way is to use the stable isotopic composition of CO (Brenninkmeijer et al., 1999; Kato et al., 1999b; Röckmann et al., 2002). ~~Only~~ Of the stratospheric observations, only Jost et al. (2004) measured CO_2, allowing $\Delta CO : \Delta CO_2$ to be quantified, confirming the smoke's origin. These source signatures have been successfully used in many ground-based and airborne studies on wildfire smoke plumes in the troposphere (e.g. Andreae et al., 2001; Bergamaschi et al., 1998; Tarasova et al., 2007).

This work presents the first balloon-borne CO and CO_2 observation of a wildfire smoke plume in the stratosphere. The Air-Core sampling technique (Karion et al., 2010; Membrive et al., 2017) provides an accurate measurement of enhancement ratios

55 of $\Delta CO : \Delta CO_2$; where the LIghtweight Stratospheric Air Sampler LISA (Hooghiem et al., 2018) is used to collect larger samples that allow for the determination of the carbon and oxygen stable isotopic compositions of CO. A back-trajectory analysis was performed using the Chemistry Lagrangian Model of the Stratosphere, CLaMS (McKenna et al., 2002), to determine the source region and ~~fire~~ injection date, which helps in the quantification of the chemical loss of CO by OH. The trace gas and isotopic composition measurements are used to confirm the wildfire smoke ~~burning origin~~ as origin of the plume. Finally

60 the observations are used to estimate the fraction of tropospheric air in the enhanced smoke plume.

2   Methods

2.1   Sampling instruments and flights

The air sampling was done with LISA (Hooghiem et al., 2018) and an AirCore (Karion et al., 2010). Both instruments are capable of sampling the stratosphere, and can be flown using small weather balloons which are easy to operate. We refer to the

65 original references for the details; here we present a brief description.

An AirCore is a long coiled thin tube, with one end open and one end closed. The AirCore passively takes an air sample during descent, relying on increasing atmospheric air pressure to push air into the tube. A magnesium perchlorate dryer is positioned at the inlet in order to dry the incoming sample. The AirCore used in this study consisted of two pieces of stainless-steel tubing with SilcoNert 1000 coating (Restec Inc.) to create a chemically inert and smooth surface. The first section was

70 a 40 m long tube with a 0.635 cm (1/4 inch) outer diameter; the second piece was a 60 m long tube, with 0.3175 cm (1/8 inch) outer diameter. Both tubes had a wall thickness of 0.0254 cm (0.01 inch). The benefit of an AirCore, when launched on a balloon, is the retrieval of an atmospheric profile. The ~~amount of sample retrieved per flight, was in total~~ AirCore used has a volume of 1.4 l ~~at standard temperature and pressure, and roughly~~ . The AirCore takes a sample passively, and due to the low pressure in the stratosphere only about 0.3 l of the sample is stratospheric. ~~The amount of sample is thus is limited~~Furthermore,

75 the AirCore has an estimated vertical resolution of 374 m at 200 hPa or 12 km altitude.

Contrary to the AirCore, LISA takes four samples actively using a small pump upstream of four sampling bags. The active sampling results in a larger amount of sample 180–800 ml per sample, thus allowing for isotope analysis. The four sampling bags are filled at a different altitude between 12–25 km during ascent (Hooghiem et al., 2018). The ascent speed is usually slower than that during descent. Sampling during ascent thus favours a higher vertical resolution, which is around 0.5 to 1 km

80 for the LISA samples.

Both LISA and the AirCore were launched together on the same balloon from the radiosonde facility of the Finnish Meteorological Institute at Sodankylä (67.37° N, 26.63° E, 179 m above mean sea level (a.m.s.l.)) using Totex TX3000 balloons. The balloons typically reach 30 km thus penetrating most of the atmospheric mass ($> 99\,\%$). Three flights with both instruments were performed, one on each day from 4 to 6 September 2017. Furthermore, the AirCore was flown without LISA on

85 7 September. ~~Additional~~ In addition to the AirCore and LISA, a Vaisala radiosonde (RS-92SGP) was added to the payload for

collocated measurements of temperature, pressure, and relative humidity ~~measurements. Furthermore, the radiosonde collects GPS and altitude information~~ as well as GPS location and altitude during flight (Dirksen et al., 2014).

The AirCore was analysed for CO_2, CH_4, and CO mole fractions, for details see Section 2.2.1. In this work only the AirCore CO_2 and CO profiles are used. LISA samples have been analysed for CO_2, CH_4, and CO mole fractions, see Section 2.2.1, and the stable isotopic composition of CO, see Section 2.2.3. Here only the LISA CO mole fraction and CO isotope measurements at the plume altitude, see Section 3.1, are used, one sample from 5 and 1 sample from 6 September 2017. Although measured, LISA CO_2 appears to suffer from a bias as concluded from comparison with AirCore measurements, see Hooghiem et al. (2018).

2.2 Measurements

2.2.1 Mole fraction measurements of CO_2 and CO

Directly after the payload was retrieved from the landing location, the AirCore and LISA samples were analysed for CO_2, CH_4, and CO mole fractions using the Cavity Ring-Down Spectroscopy (CRDS) technique~~(analyser model~~. Two analysers were used, to allow for simultaneous analyses of both AirCore and LISA samples after payload retrieval (analyser models used: Picarro G2401 (AirCore), Picarro G2401-m (LISA) see e.g. (Crosson, 2008) for more information on the CRDS-analyser). A calibration gas was also measured to link the mole fraction measurements to the following World Meteorological Organization, or WMO, scales: X2007 (CO_2) (Zhao and Tans, 2006), X2004 (CH_4) (Dlugokencky et al., 2005) and X2014A (CO). Throughout this work, the abbreviation ppm is used for $\mu mol\,mol^{-1}$, and ppb for $nmol\,mol^{-1}$.

After post-processing of the analyser output, dry mole fractions were obtained using the instrument specific but well-defined analyser response to H_2O (Rella et al., 2013; Chen et al., 2010). This is especially relevant for the LISA samples (Hooghiem et al., 2018); the range of water vapour mole fractions was 0.03–0.15 %, partly because of diffusion into the sampling bag (Hooghiem et al., 2018). The mole fraction results of the AirCore were further processed to give vertical profiles as described in Karion et al. (2010); Membrive et al. (2017). The uncertainty of the AirCore measurements typically is 0.1 ppm for CO_2, 2 ppb for CH_4 and 2 ppb for CO. Measurements on samples collected with the LISA sampler have an uncertainty of 0.14 ppm for CO_2, 2.3 ppb for CH_4 and 7.8 pbb for CO. This uncertainty includes analyser precision, calibration transfer, a dead volume bias, and storage bias; for the technical details we refer to Hooghiem et al. (2018).

2.2.2 LISA sample transfer and storage

As described in Hooghiem et al. (2018), the bags used in the LISA sampler provide limited stability to the sample. Therefore, the samples were transferred into 350 ml glass flasks, after the CRDS analysis. The flasks have a Rotulex connection and are sealed with two Viton-70 O-rings, providing better sample stability than single O-ring configuration (Sturm et al., 2004). The flasks and transfer lines were evacuated using a vacuum pump (flasks were evacuated using an Adixen Drytel 1025, the transfer lines using a Vacuubrand MD 1) before the samples were introduced into the flasks. As the bags are compressible, the sample was pushed into the flask until local ambient atmospheric pressure was reached, typically 950 hPa, or until the sample was

fully expanded into the glass container at a pressure lower than ambient pressure. The air samples were stored in these glass flasks until they were analysed in the laboratory for the isotopic compositions of CO. The storage ~~period was 3–7 months.~~ for the data presented in this work were 55 days for the background sample and 84 for the plume days, see below in Table 4. The drift in CO mole fractions is estimated to be 0.05 ppb day^{-1}, from a storage test of stratospheric samples (20–30 ppb initially, average increase of 36 ppb over 2 years, not reaching ambient mole fractions). The stability of the stable isotopic composition was not assessed directly, but based on an estimate using the drift in mole fractions, it was estimated to be small in general. Yet, it can not be excluded that the isotopic measurements, see below, are biased by more than a ‰.

2.2.3 Analysis of stable isotopic composition of CO

The LISA samples were shipped to the Institute for Marine and Atmospheric research Utrecht (IMAU) for analysis of δ^{13}C and δ^{18}O in CO. The samples were analysed using Continuous-Flow Isotope-Ratio Mass Spectrometry (CF-IRMS) (Pathirana et al., 2015). The δ ~~values~~ values for carbon are reported on the Vienna Peedee Belemnite scale (VPDB), whereas oxygen values are reported on the Vienna Standard Mean Ocean Water scale (VSMOW). For details about the measurements we refer the reader to Pathirana et al. (2015). Briefly, the sample is carried, using He as a carrier gas, through an Ascarite and magnesium perchlorate trap, removing CO_2 and H_2O from the sample. N_2O and any remaining CO_2 are removed by means of a cryogenic trapping using liquid N_2. Then the CO is converted to CO_2 with the aid of the Schütze reagent. A second cryogenic trap isolates the CO_2 derived from CO, which, after ~~passing through~~ purification on a GC column, is fed to an IRMS via an open split system. The IRMS analyses δ^{13}C and δ^{18}O. The δ^{18}O is corrected for the additional oxygen atom added in the conversion to CO_2 as described by (Pathirana et al., 2015). The CF-IRMS system in this study was the same as that described in (Pathirana et al., 2015), with one exception. The CF-IRMS analysis used about 150 ml sample, and requires a sufficiently high upstream pressure (> 900 hPa absolute) to maintain a constant sample flow. As mentioned in Section 2.2.2, the starting pressure of the LISA samples was equal to or lower than 950 hPa, and it decreased rapidly during a measurement due to the small flask volume of 350 ml. Therefore, the pressure in the flasks was increased during the measurement using CO-free synthetic air. ~~Each sample was measured at least twicewhile being diluted in this way~~Samples were measured twice, where the first measurement was performed without dilution.

As the samples were measured at very low ~~concentrations~~mole fractions, meaning very low peak areas in the IRMS measurement, special attention was paid to ~~estimate the~~ quantify the potential effect of non-linearity on the reported isotopic composition. Dilution tests showed detectable non-linear behaviour for δ^{18}O and δ^{13}C below a peak area corresponding to mole fractions of approximately 10 ppb and 15 ppb respectively, but none of the samples presented here were measured at such low peak areas. Therefore we can consider the non-linearity effect negligible.

The average analytical precision for this dataset, estimated from the reproducibility of repeated sample measurements, was 0.5 ‰ for δ^{13}C and 0.5 ‰ for δ^{18}O.

2.3 Characterisation of the plume

2.3.1 Back-trajectory analysis

To determine the origin of the observed air masses with enhanced CO and CO_2 mole fractions (see Section 3.1), back-trajectories were calculated with the trajectory module of the Chemical Lagrangian Model of the Stratosphere (CLaMS) (McKenna et al., 2002) driven by ERA-Interim meteorological data with a resolution of $1°$ by $1°$ (Dee et al., 2011). The mixing and advection schemes of CLaMS are capable of resolving the filamentary structures that exist in the stratified stratosphere. The CLaMS trajectories are calculated on isentropic surfaces with a 30 min time step. The vertical displacement from the isentropic surfaces calculated from diabatic vertical velocities (Ploeger et al., 2010).

Khaykin et al. (2018) showed increased vertical transport of wildfire smoke due to heating induced by aerosols. As this additional vertical velocity component is not included in the model computation, it is difficult to directly backtrack the air masses by a single trajectory. Therefore, a correction for the vertical displacement is determined in correspondence with CALIOP elastic backscatter-ratios at 532 nm. The back trajectories started from the observed CO maximum $p = 155$ hPa, 13.6 km. In the present study, only the advection scheme is used. First, information on the aforementioned additional vertical motion is lacking. Secondly, only synoptic scale transport is of interest to determine the source region of the smoke.

2.3.2 Enhancement ratio of CO to CO_2 in the plume

CO and CO_2 are co-produced in burning processes, and their emissions into the atmosphere result in an enhancements, ΔCO and ΔCO_2, compared to background air. The enhancement ratio of $\Delta CO : \Delta CO_2$ is typically high for wildfires, and decreases over time due to photochemical loss of CO (Mauzerall et al., 1998). The enhancement ratio is conserved during mixing, if the background is constant. Thus, in this case, the enhancement ratio can be directly obtained as the slope of a linear regression performed on AirCore CO and CO_2 data for the two separate AirCore flights that sampled the plume on 4 and 5 September 2017.

Alternatively, measurements of background air can be used to compute the enhancement ratio directly. It is assumed that the flight performed on 6 September is representative for the background air adjacent to the plume. CO and CO_2 data are smoothed using a moving average with an averaging window of 25 data points in order to reduce the analyser noise in CO. The background air is interpolated on isentropic surfaces to the observed plume altitude, and the enhancements are computed directly on isentropic surfaces. As we will see, the total enhancement in CO_2 is small, and only the results with $x(CO_2) > 0.2$ ppm are used in the subsequent computation of the enhancement ratios.

2.3.3 Determination of the source signature

Table 1. Four main sources of CO, and their isotopic source signatures. [a](Stevens et al., 1972), [b](Brenninkmeijer, 1993), [c](Stevens and Wagner, 1989), [d](Bergamaschi et al., 1998), [e](Saurer et al., 2009), [f](Manning et al., 1997), [g](Brenninkmeijer and Röckmann, 1997), and [h](Vimont et al., 2019). Table based on the most recent compilation of source signatures by Vimont et al. (2019).

Source	^{13}C (VPDB)	Uncertainty	^{18}O (VSMOW)	Uncertainty
Fossil fuel combustion[a,b]	$-27.5\,‰$	$\leq 1\,‰$	$23.5\,‰$	$\leq 1\,‰$
Biomass burning[c,d,e,f]	-12–$-25\,‰$	1–$3\,‰$	10–$18\,‰$	1–$3\,‰$
CH$_4$ oxidation[f,g]	$-52.6\,‰$	1–$3\,‰$	$0\,‰$	$> 3\,‰$
NMHC oxidation[c,g,h]	$-32\,‰$	1–$3\,‰$	0–$4\,‰$	$> 3\,‰$

180 Various sources of CO emit CO with different isotopic composition, see Table 1, after Vimont et al. (2019). The isotopic composition is then modified by physical (mixing) and chemical (oxidation) processes in the atmosphere. Inversely, knowledge of these processes can be used to determine the source of an observed anomaly in isotopic composition of a tracer in an airmass. Here, two methods are outlined to determine the source signature. The first method describes a simple mixing process, without oxidation. The second method describes the determination of the source signature with both mixing and oxidation.

185 The method usually employed to determine the source signature of ~~a pollutionsource~~ an observed pollution, is to assume that a measured air parcel is the result of mixing of background air and a polluting source, i.e. two end-member mixing. Then the following mass balance applies to the mole fractions x:

$$x_{\text{ap}} = f_{\text{bg}}x_{\text{bg}} + f_{\text{src}}x_{\text{src}} \tag{1}$$

where ap denotes the air parcel, bg means background, and src means the pollution source. ~~A similar~~ Here f_{bg} is the fraction
190 of molecules of bg in ap, and similar for src; mass conservation requires $f_{\text{bg}} + f_{\text{src}} = 1$. A similar mass balance can be written for the stable isotopes~~-~~:

$$x_{\text{ap}}\delta^{13}\text{C}_{\text{ap}} \overset{=}{\approx} f_{\text{bg}}x_{\text{bg}}\delta^{13}\text{C}_{\text{bg}} + f_{\text{src}}x_{\text{src}}\delta^{13}\text{C}_{\text{src}} \tag{2}$$

~~Equation~~since there is a small loss of tracer (Tans, 1980), hence the $\approx$ sign instead of $=$. Equation 1 and Equation 2 can be solved to yield:

195 $$\delta^{13}\text{C}_{\text{ap}} = \left(\delta^{13}\text{C}_{\text{bg}} - \delta^{13}\text{C}_{\text{src}}\right)\frac{f_{\text{bg}}x_{\text{bg}}}{x_{\text{ap}}} + \delta^{13}\text{C}_{\text{src}} \tag{3}$$

~~Which~~ which results in a linear relation between $\delta^{13}\text{C}_{\text{ap}}$ and x_{ap}^{-1} if $(\delta^{13}\text{C}_{\text{bg}} - \delta^{13}\text{C}_{\text{src}})f_{\text{bg}}x_{\text{bg}}$ is assumed to be constant. This relation was first recognized by Keeling (1958) and is a special case of the more general Miller-Tans method (Miller and Tans, 2003). A relation for ^{18}O, equivalent to Equation 3, can be derived as well.

~~This relation~~ Two end-member mixing, as described above, can be safely applied to CO if it can be assumed that removal by
200 OH is negligibly small. Recent applications of this methods can be found in e.g. Vimont et al. (2019, 2017); Gromov and Brenninkmeijer (2

Table 2. The coefficients a, b, and c for Equation 7 for ^{13}C and ^{18}O.

Isotope	a	b	c
^{13}C	-0.00655	0.02269	0.00947
^{18}O	-0.01191	0.00603	-0.00341

However, the age of the observed air parcel, 24–25 days (see Section 3.2), is too old to ignore the chemical reaction of $CO + OH$. Therefore the evolution of CO in the plume is modelled as follows:

$$\frac{dn(CO)}{dt} = -k_{er}\left(n(CO) - n_{bg}(CO)\right) - k_1 n(OH)n(CO) \tag{4}$$

where k_{er} is the entrainment rate in ~~s^{-1}~~ s^{-1}. $n(X)$ is the number density of species X, in ~~cm^{-3}~~ cm^{-3}. The reaction rate of
205 the reaction $CO + OH$, k_1, was taken from McCabe et al. (2001):

$$k_1 = 1.57 \cdot 10^{-13} + 3.54 \cdot 10^{-33} n \tag{5}$$

in cm^3 s^{-1}, where n is the number density of air in cm^{-3}. The number density of OH is taken to be $2.7 \cdot 10^6$ cm^{-3} between 0–4 km altitude, $1.6 \cdot 10^6$ cm^{-3} between 4–8 km altitude and $1.2 \cdot 10^6$ cm^{-3} for altitudes >8 km (Mauzerall et al., 1998), although Mauzerall et al. (1998) specifies this value only for the range 8–12 km. ~~The production of from NMHCs and is~~
210 ~~assumed to be insignificant, and therefore neglected.~~

After separating the variables of Equation 4, integrating, and solving for $n(CO)(t)$ using the boundary condition that at $n(CO)(t=0) = n_0(CO)$, this yields (see Section A for the derivation):

$$n(CO)(t) = \left(n_0(CO) + \frac{k_{er}n_{bg}(CO)}{-k_{er} - k_1 n(OH)}\right)\exp\left(-\left(k_{er} + k_1 n(OH)\right)t\right) - \frac{k_{er}n_{bg}(CO)}{-k_{er} - k_1 n(OH)} \tag{6}$$

An equivalent equation to Equation 6 can be written for ^{13}CO and C^{18}O. The reaction rates for these minor isotopologues
215 can be determined from fractionation factors:

$$\alpha = \frac{k_{minor}}{k_{major}} = \frac{1}{a + bp + cp^2} \tag{7}$$

where p is the atmospheric pressure in bar. The coefficients a, b, and c are fit to combined datasets from Röckmann et al. (1998); Stevens et al. (1980); Smit et al. (1982) and are obtained from Gromov (2013), and can be found in Table 2. k_{minor} and k_{major} are the reaction rates of the rare and abundant isotopologues respectively. We assume that $k_{major} = k_1$ as in Equation 5.
220 Now k_{er} can be written as (see Section A for the derivation):

$$k_{er} = \frac{-\ln\left(1 - f_{strat}\right)}{t} \tag{8}$$

8

If the age, t, of the air parcel is known, then k_{er} can be evaluated as a function of f_{strat} which is the total fraction of stratospheric air entrained in the air parcel.

In principle f_{strat} is unknown, but it can be evaluated over the full range from 0 to 1 to give ~~an idea what range of source~~ ~~signatures is. Then the source signature can be estimated, and~~ provide a range of possible source signatures. Then this range can be compared to the ~~known sources and their signatures~~ signatures of possible sources, see Table 1. Furthermore, a best estimate can be made using f_{strat} found from Section 2.3.4.

~~Temperature~~ Initial temperature and pressure are known from the observation made by LISA; temperature and pressure are assumed to be constant during transport. Then Equation ~~4~~ 6 can be used to obtain the CO, ^{13}CO, and C^{18}O number densities from which δ^{13}C and δ^{18}O can be determined ~~at an earlier time compared to the observation~~ throughout the simulation of the air parcel in the stratosphere.

~~Four main sources of CO, and their isotopic source signatures. ^{a}(Stevens et al., 1972), ^{b}(Brenninkmeijer, 1993), ^{c}(Stevens and Wagner, 19~~ ~~, ^{d}(Bergamaschi et al., 1998), ^{e}(Saurer et al., 2009), ^{f}(Manning et al., 1997), ^{g}(Brenninkmeijer and Röckmann, 1997), and ^{h}(Vimont et al., 2~~ ~~. Table based on the most recent compilation of source signatures by Vimont et al. (2019). Source ^{13}(VPDB) Uncertainty~~ ~~^{18}(VSMOW) Uncertainty Fossil fuel combustion^{a,b} $-27.5\text{‰}\leq 1\text{‰}23.5\text{‰}\leq 1\text{‰}$Biomass burning^{c,d,e,f} -12−$-25\text{‰}1-3\text{‰}10-18\text{‰}1-3\text{‰}$~~ ~~oxidation^{f,g} $-52.6\text{‰}1-3\text{‰}0\text{‰}>3\text{‰}$~~ ~~NMHC oxidation^{c,g,h} $-32\text{‰}1-3\text{‰}0-4\text{‰}>3\text{‰}$~~

2.3.4 ~~Stratosphere-troposphere exchange estimate~~ Estimate of the tropospheric air fraction based on the in situ observations

~~By combining the isotope and mole fraction measurements in a simple mixing model, the contribution of tropospheric and stratospheric~~ ~~air to the observed plume can be quantified using 3 end-member mixing~~ The mass balances in Equations 1 and 2 can be extended to allow for more than two end members. Assuming mixing of stratospheric air, tropospheric air, and wildfire smoke, the mass balance for the mole fraction of the ~~air parcel, called ap (air parcel)~~ observed plume, here called the air parcel or ap, sampled by LISA is ~~can be written~~ as follows:

$$x_{\text{ap}} = f_{\text{w}}x_{\text{w}} + f_{\text{t}}x_{\text{t}} + f_{\text{s}}x_{\text{s}} \tag{9}$$

and for the stable isotopes, where the same approximation in Equation 2, discussed in Section 2.3.3, is used:

$$x_{\text{ap}}\delta^{13}\text{C}_{\text{o}} = f_{\text{w}}x_{\text{w}}\delta^{13}\text{C}_{\text{w}} + f_{\text{t}}x_{\text{t}}\delta^{13}\text{C}_{\text{t}} + f_{\text{s}}x_{\text{s}}\delta^{13}\text{C}_{\text{s}} \tag{10}$$

and

$$x_{\text{ap}}\delta^{18}\text{O}_{\text{ap}} = f_{\text{w}}x_{\text{w}}\delta^{18}\text{O}_{\text{w}} + f_{\text{t}}x_{\text{t}}\delta^{18}\text{O}_{\text{t}} + f_{\text{s}}x_{\text{s}}\delta^{18}\text{O}_{\text{s}} \tag{11}$$

Here f and x ~~denote volumetric air fraction and mole fraction respectively~~ are defined in Section 2.3.3. Subscripts w, t, and s denote the "wildfire smoke", "tropospheric", and "stratospheric" end-members respectively, whereas ap denotes "air parcel".

An end-member is defined here by its carbon and oxygen isotopic composition, and by its mole fraction, and it is possible to distinguish the air parcel from others based on those parameters. Mass balance requires:

$$f_w + f_t + f_s = 1 \tag{12}$$

This is an extension of two end-member mixing presented by Equation 1 and Equation 2. Combining Equation 9, Equation 12, Equation 10, and Equation 11 yields a system of four linear equations:

$$
\begin{bmatrix}
1 & 1 & 1 \\
x_w & x_t & x_s \\
x_w\delta^{13}c_w & x_t\delta^{13}C_t & x_s\delta^{13}C_s \\
x_w\delta^{18}O_w & x_t\delta^{18}O_t & x_s\delta^{18}O_s
\end{bmatrix}
\begin{bmatrix}
f_w \\
f_t \\
f_s
\end{bmatrix}
=
\begin{bmatrix}
1 \\
x_{ap} \\
x_{ap}\delta^{13}C_{ap} \\
x_{ap}\delta^{18}O_{ap}
\end{bmatrix}
+
\begin{bmatrix}
M_r \\
x_r \\
x_r\delta^{13}C_r \\
x_r\delta^{18}O_r
\end{bmatrix}
\tag{13}
$$

The second term on the right-hand side, $[M_r, x_r, x_r\delta^{13}C_r, x_r\delta^{18}O_r]$, is the residual vector that ensures equality of the over-constrained problem, with M_r denoting the normalized mass. Note that ideally this term is zero.

This set of equations is normalized to the observations and weighted so that all residuals, are of equal importance, except for the mass balance. The mass balance was given extra weight, which conforms to the assumption that the observed air parcel is purely a mixture of tropospheric, stratospheric, and wildfire smoke.

Equation 13 was solved for f_w, f_t, and f_s by minimizing the residuals $[M_r, c_r, c_r\delta^{13}C_r, c_r\delta^{18}O_r]$ using a non-negative least square algorithm (Lawson and Hanson, 1995).

The values assumed for the end-members variables are presented in Table 3, and rely on different sources and will be discussed in detail below. All of the end-members have a different uncertainty. In order to capture the effect of these uncertainties, a Monte-Carlo simulation was performed, rather than a single calculation, to estimate f_w, f_t, and f_s. The end-member values are randomly drawn from the respective distributions assumed, also presented in Table 3.

The stratospheric end-member definition and the plume observation are based on the balloon-borne observations by LISA, outside and inside the plume, respectively, see Table 4. The plume mole fractions are assigned an uncertainty equal to the measurement uncertainty of LISA. The uncertainty of the stratospheric end-member is also a good measure of the variability surrounding the plume based on AirCore measurements between 12 and 14.5 km of 6 and 7 September. The uncertainty of the isotopic composition in the Monte-Carlo simulation is set equal to twice the measurement uncertainty (see Section 2.2.3) because the stratospheric variability of the stable isotopic composition of CO is unknown. Secondly, there may have been a small drift in the isotopic composition.

In an additional Monte-Carlo experiment, the plume observation is corrected for removal by OH, using Equation 6 with $k_{er} = 0$. t is equal to the 25 days of transport. This definition of the plume observation appears in Table 3 as "ap OH-corrected". A number density OH of $1.2 \cdot 10^6$ cm^{-3} is used in the calculation.

~~The end-member definitions used are presented in Table 3. The stratospheric end-member definition and the plume observation are based on the balloon-borne observations presented in this paper, and are well defined. The mole fractions are assigned an uncertainty equal to the measurement uncertainty, which, considering the observations from AirCore, is also a reasonable~~

Table 3. End-member ~~definitions~~ variables and air parcel values and uncertainties as used in the Monte-Carlo simulation. Mole fractions are given in ppb, and δ ~~-values~~ values in ‰. ~~The large range for the plume mole fraction does not significantly affect the results.~~ Here ap means air parcel, and is based on the in situ ~~enhancement~~ observation made by LISA, see Table 4. ~~An additional Monte-Carlo run~~ The uncertainties for the mole fractions is ~~made with~~ set equal to the uncertainty of the LISA observations as discussed in Section 2.2.1. The uncertainty of the isotopic composition of the stratosphere, ap, and ap OH-corrected are set equal to twice the measurement uncertainty, see text for a discussion. In the ap OH-Corrected ap is corrected for removal by OH ~~using~~ using Equation 6 with $k_{\mathrm{er}} = 0$. A number density OH of $1.2 \cdot 10^6$ cm^{-3} is used. The stratospheric end-member definition is based on the in situ observation of stratospheric background presented in Table 4. The tropospheric and wildfire smoke end-member definitions are based on observations from earlier work, see text. The large range for the plume mole fraction does not significantly affect the results of the Monte-Carlo simulation.

Airmass	Variable	Mean	1-σ	Distribution
Wildfire smoke	CO	$0.5 \cdot 10^6$ to $1.5 \cdot 10^6$	-	uniform
	δ^{13}C	-24.4 to -21.3	-	uniform
	δ^{18}O	16.3 to 18.0	-	uniform
Stratosphere	CO	34	8	normal
	δ^{13}C	~~-1.0~~ -29.6	1.0	normal
	δ^{18}O	~~-29.6~~ -1.0	1.0	normal
Troposphere	CO	72	8	normal
	δ^{13}C	-32 to -28	-	uniform
	δ^{18}O	-4 to 0	-	uniform
ap	CO	74	8	normal
	δ^{13}C	-28.8	1.0	normal
	δ^{18}O	4.3	1.0	normal
ap OH-Corrected	CO	~~111.0~~ 116.0	8	normal
	δ^{13}C	~~-26.2~~ -27.4	1.0	normal
	δ^{18}O	~~9.2~~ 9.4	1.0	normal

~~measure of natural variability. Yet, for the stable isotopes a larger uncertainty was assumed, twice the measurement uncertainty, to account for unknown variability of the stratospheric background and variability within the plume.~~

The wildfire smoke and tropospheric end-member rely on measurements from earlier publications, which introduces a large uncertainty on end-member definitions. First of all, the isotopic composition and mole fraction of tropospheric CO exhibits temporal, both seasonal and annual, and latitudinal gradients (Bergamaschi et al., 2001; Mak et al., 2003), which complicates the characterisation of the tropospheric end-member. The smoke source region lies between $65°$–$75°$ N (see Section 3.2), with south-westerly surface winds coming from the Pacific Ocean (Peterson et al., 2018). Tropospheric CO mole fractions used are based on measurements of CO from Midway island (Petron et al., 2019), and reported to be typically 72 ± 8 ppb in the six weeks preceding the event. The range of isotopic-composition values considered are therefore obtained from measurements made at Izana ($28°$ N and $16°$ W, which is representative for CO in air that travels over the ocean at mid-latitudes. δ^{13}C ranges between -32 and -28 ‰ and δ^{18}O ranges between -4 and 0 ‰ (Bergamaschi et al., 2001; Mak et al., 2003). The mole fractions obtained here from Midway island are consistent with those co-reported with the δ^{13}C and δ^{18}O ranges by Bergamaschi et al. (2001); Mak et al. (2003).

In addition, the wildfire smoke signature, see Table 1, is subjected to a large variability due to the type of burned plants (categorised as C3 or C4, with a difference in photosynthesis), the burning temperature, and possibly the groundwater isotopic composition (Kato et al., 1999b). Since the wildfire originated in Canada, the fuel consisted mainly of C3 plants which are typically more depleted in δ^{13}C. Furthermore, the fire was energetic enough to trigger a pyro-Cb event, which makes it reasonable to assume that it was in an efficient burning regime, which typically leads to higher δ^{18}O. Therefore, the range of isotopic composition of CO of the wildfire smoke assumed here is according to atmospheric measurements around forest fires (Brenninkmeijer et al., 1999), see Table 3. This is thus a subset from the range of signatures displayed in Table 1, which provide a more general summary.

The mole fraction of wildfire smoke is also an unknown, but it is clear that it is far larger than both the stratospheric and tropospheric background mole fraction, thus a large mole fraction ensures, by virtue of Equation 12, $f_\mathrm{w} \ll f_\mathrm{s}$ and $f_\mathrm{w} \ll f_\mathrm{t}$.

~~In order to capture the above-mentioned uncertainty, a Monte-Carlo simulation was performed. The variable input parameters are randomly drawn from the respective distributions assumed, also presented in Table 3.~~

Since the mole fraction and isotopic composition of the smoke plume and the tropospheric end-members are ill-defined, partially due to natural variability, the Monte-Carlo results are filtered. Solutions to Equation 13 are only allowed if the residuals are smaller than the measurement 1-sigma uncertainty attributed to our stratospheric observations, e.g. 8 ppb for the mole fractions and 0.5 ‰ for the δ-values. A solution is thus only allowed when it is consistent with the observations and the reported isotopic composition range published in literature. Furthermore, the solution requires all fractions to be larger than 0, where four significant figures were considered, to avoid a large amount of unrealistic solutions.

3 Results

3.1 Observation of a stratospheric CO enhancement

The CO measurements of both the LISA sampler and the AirCore are presented in Figure 1a. A clear carbon monoxide enhancement was observed between 13 and 14 km altitude on 4 and 5 September 2017. The plume, extending over roughly 1 km in altitude, was well above the tropopause which can be seen from the CO gradient below 13 km. The tropopause height was determined based on the lapse rate from the radiosonde temperature measurements to be 12 km on both flights, confirming that the observed plume is above the tropopause. The potential temperature was $\theta \approx 350$ K at 13 km and $\theta \approx 380$ K at 14 km, which classifies this part of the stratosphere as the extra-tropical lowermost stratosphere (Holton et al., 1995). The observed CO_2 mole fraction Figure 1b showed a slight increase in the same layer, which allowed for determination of the enhancement ratio, $\Delta CO : \Delta CO_2$ see Section 3.3.
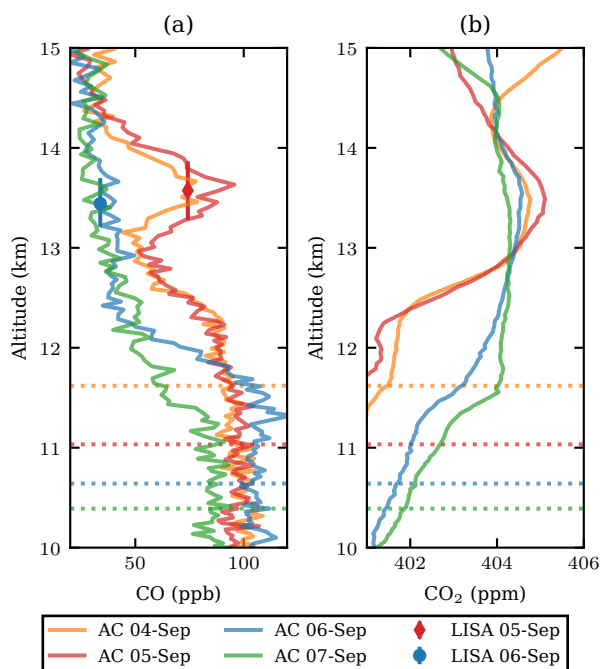
Figure 1. The CO, (a), and CO_2, (b), profiles from AirCore (lines, abbreviated as AC in the legend) and the LISA sampler (markers). The profiles are shown between 12 and 15 km altitude and are coloured by date. The LISA sampler vertical error bars represent the total vertical coverage of the sample, with the mean altitude as shown. For mole fraction measurement uncertainties, the reader is referred back to Section 2.2.1. The tropopause is indicated as a dashed line, matching the colour of the respective AirCore flight.

3.2 Origin and age of the plume based on back-trajectory analysis

325 ~~Figure~~ A first set of 6 back-trajectories was initialized from the altitude of the observed CO-peak maximum ($p = 155$ hPa, 13.6 km) and the altitudes where the CO-enhancement was half of the maximum (166 and 148 hPa, 13.3–13.8 km see Figure 1) starting on 4 and 5 September. After that, matches with CALIOP night-time observations were determined as follows. For each orbit time, the minimum distance of the trajectory and the orbit location was calculated. Wherever the distance was smaller than 250 km, the observed backscatter ratio was investigated for nearby atypical aerosol enhancements. In this way the smoke

330 could be traced back to the injection date and region.

Figure 2a shows the CALIOP backscatter ratio at 532 nm (R_{532}) as well as the location ~~the of~~ of the back-trajectory result on 3 September, with a match distance of 47 km for the lower starting altitude of the trajectories starting on 4 September. The match distance of the trajectories from the centre and upper altitude are above 250 km (282 and 434 km) and are not shown here. To track the aerosol cloud, new back-trajectories were initialized starting from 3 September exactly where aerosol

335 cloud is observed in the CALIOP data (Figure 2b white crosses). Other matches between back-trajectory location and CALIOP overpass was always at a layer of aerosol enhancement and no new back-trajectories were initialized. These match results are not shown.

Again matches with CALIOP orbits were calculated. Figure 2c and d are similar to Figure 2a and b but for 20 August, where a difference in altitude is observed. This altitude mismatch might be due to the additional vertical motion related to additional

340 radiative heating of the smoke plume, which is absent in the CLaMS model. Similarly to 3 September, in a third step, back-trajectories were calculated from this observed aerosol cloud. On 14 August, the back tracked air parcels are within the range of the aerosol cloud that was observed by the Ozone Mapping Profiler Suite (OMPS), as shown by Peterson et al. (2018). Figure 3a shows the computed trajectories and the locations where re-initialization was performed. The result of the back-trajectory analysis is shown to match the location of the aerosol enhancement observed by OMPS and CALIOP in Figure~~3~~ 3b.

345 Figure 2f shows the location on 12 August, before the injection. This cloud is located over British Columbia and was caused by the pyro-convection as concluded by Peterson et al. (2018). Further back in time, the CALIOP backscatter data do not show any aerosol enhancement where a location match between CALIOP and the back-trajectory was found. Therefore, the origin of the observed plume could be confirmed by this piecewise trajectory analysis that accounts for the vertical transport due to heating caused by the fire. Furthermore, the age of the plume at the time of observation was 24–25 days.

350 ## 3.3 Enhancement ratios of CO/CO$_2$ of the plume

Figure 4 shows a scatter plot of the observed CO and CO$_2$ mole fractions in the plume, both measured CO and CO$_2$ data and CO corrected for removal by OH (see below). The enhancement ratio on 4 September, 40 ± 2 ppb ppm^{-1}, is higher than that on 5 September, 34 ± 1 ppb ppm^{-1}. The mole fraction enhancement ratio, 34–40 ppb ppm^{-1}, falls in the range of fresh to aged biomass burning plumes, as defined by Mauzerall et al. (1998). The regression coefficients are high ($r^2 > 0.8$).

355 ~~vs scatter plot at the observed enhancement in the AirCore profiles from 4 and 5 September 2017. The and data presented here lies between the pressure levels 167 and 140 (13.2–14.3) on 4 September and between 175 and 140 (12.9–14.4) on 5 September. The -corrected versus is also plotted.~~
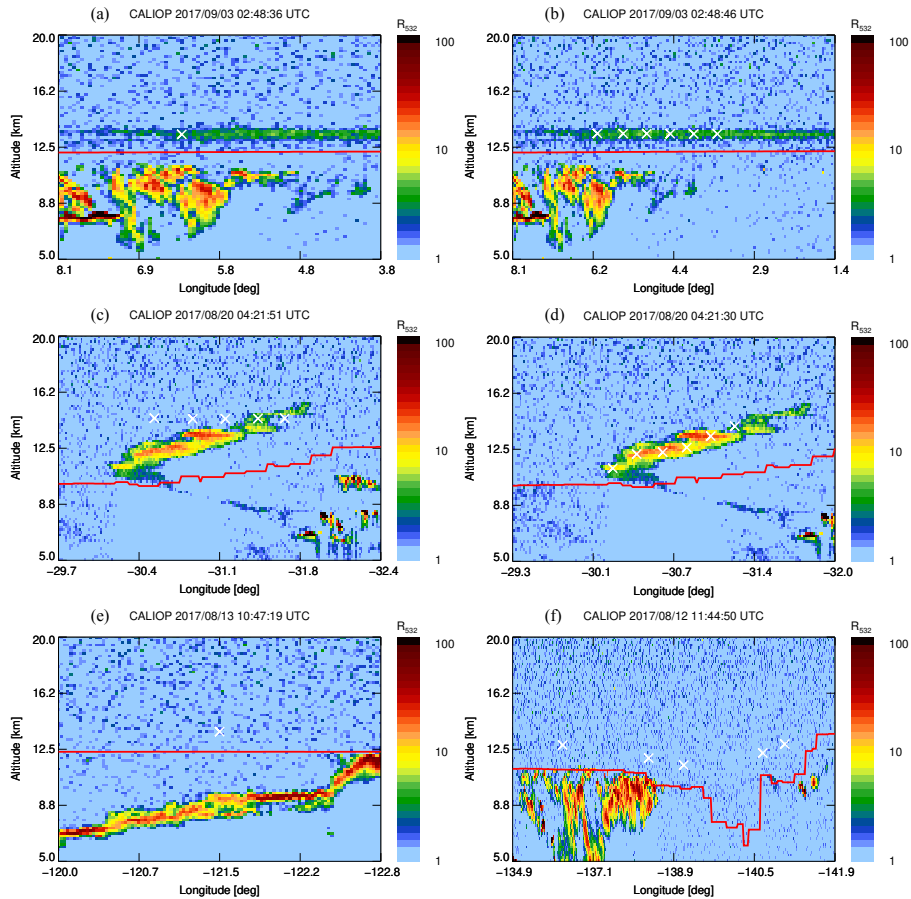
14

Figure 2. Altitude-Longitude plot with CALIOP backscatter ratio and back-trajectory results. The red line indicates the thermal tropopause from CALIOP. a) ~~a~~ A match on 3 September between CALIOP and the lower CLaMS ~~back-trajectories~~ back-trajectory starting on ~~5~~4 September, indicated by the white cross. The distance between the back-trajectory result and the plotted white cross is 47 km. b) newly initialized back-trajectories on 3 September, white crosses. c) ~~the~~ The matches on 20 August, between CALIOP backscatter and the results from the back-trajectories initialized on 3 September, where the distance between back-trajectory result and CALIOP overpass is smaller than 250 km. Note the altitude discrepancy ~~. During~~ showing that the ~~period between 20 August and 3 September~~ vertical displacement is not precisely reproduced in the ~~correspondence was always sufficiently good~~ model. d) ~~newly~~ Newly initialized back-trajectories on 20 August. e) Match location on 13 August with a vertical distance of 3 km to the aerosol plume. f) Match location on 12 August with no clear correspondence between the trajectory and the enhanced backscatter from CALIOP~~, the~~ . The time of the location match is well before the fires, 21:00 UTC (Peterson et al., 2018).
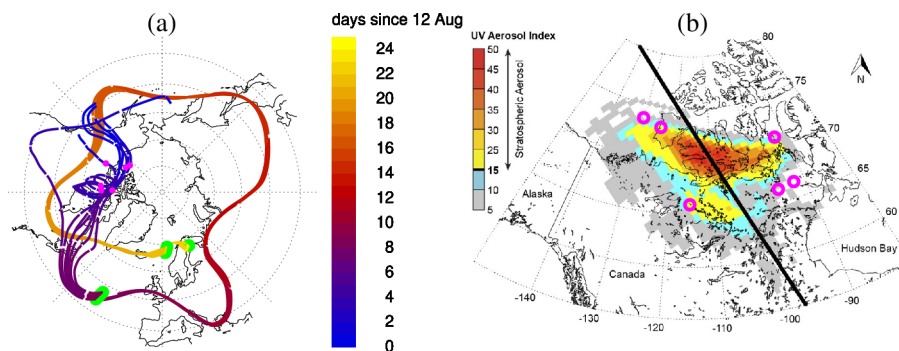
Figure 3. The panel on the left, (a), shows the computed trajectories. The three green areas show the location of the initialization and the two locations where the back-trajectories are corrected, mainly in altitude, using a CALIOP match. The latter two correspond to the panels a/b and c/d in Figure 2. The colour gradient shows days elapsed since 12 August. The pink circles in the figure on the left are showing the same location as in the figure on the right, (b). On the right, (b), a reprint of the figure from (Peterson et al., 2018, Fig. 3) showing the ultra-violet aerosol index from the Ozone Mapping Profile Suite (OMPS) on 14 August with the CALIPSO satellite track in black. Here the result of the back-trajectory on 14 August are added to the figure, ~~blue~~ pink circles, coinciding with the stratospheric smoke plume. The original figure was published under a Creative Commons Attribution 4.0 International License, http://creativecommons.org/licenses/by/4.0/.

The chemical lifetime of CO against removal by OH is about 50 days in the stratosphere (e.g. Mauzerall et al., 1998). Since the age of the plume after the injection date of 12 August 2017 (Peterson et al., 2018) was 24 and 25 days for the observation on 4 September and 5 September 2017, respectively, the mole fraction of CO in the plume was thus significantly affected by chemical loss due to the reaction of CO with OH. Therefore, the observed CO mole fractions are corrected by assuming a continuous removal by OH, an approach similar to the one used by Andreae et al. (2001), using the parameters presented by Mauzerall et al. (1998).

~~The plume was transported dominantly in the stratosphere as shown by the back-trajectory analysis in Section 3.2. (Peterson et al., 2018)~~ Peterson et al. (2018) estimated the time of ~~the~~ up-draft in the troposphere to be ~~around~~ 5 hours, thus the plume was transported dominantly in the stratosphere. The dilution of the plume due to mixing with ambient air is ignored in the calculation. The OH-corrected enhancement ratios are plotted in Figure 4. The corrected enhancement ratios are 53 ± 2 and 62 ± 3 $\mathrm{ppb\,ppm^{-1}}$ ~~respectively. Thus~~ for the two days respectively. Using the uncorrected enhancement ratios, the loss of CO was estimated about 35–37 %.

Using the data obtained from the 6 September flight as background, the following enhancement ratio's are calculated: 45 ± 1 $\mathrm{ppb\,ppm^{-1}}$, uncorrected for OH, and 177 ± 2 $\mathrm{ppb\,ppm^{-1}}$, using a correction for removal by OH for 4 September (number of data points $N = 5$); 91 ± 29 $\mathrm{ppb\,ppm^{-1}}$, uncorrected for OH, and 201 ± 56 $\mathrm{ppb\,ppm^{-1}}$, using a correction for removal by OH for 5 September ($N = 50$). These are higher than those obtained from the linear fit, which is likely caused by relatively large uncertainties of the small ΔCO_2 values.
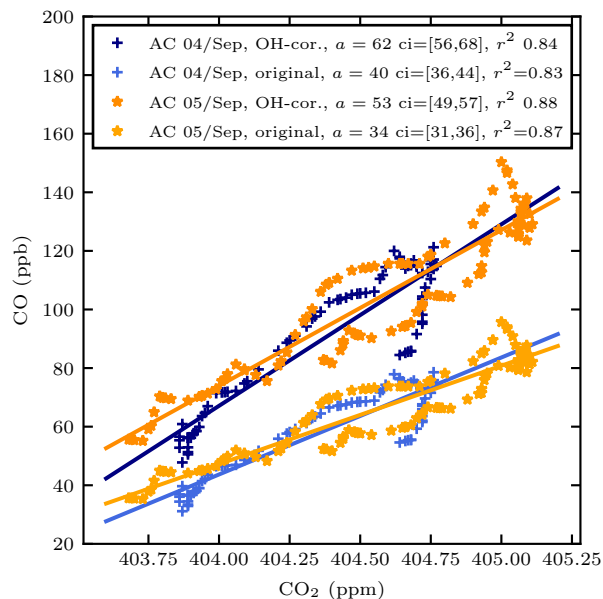
Figure 4. CO vs CO$_2$ scatter plot at the observed CO enhancement in the AirCore (AC) profiles from 4 and 5 September 2017. The data shown are between the pressure levels 167 hPa and 140 hPa (13.2–14.3 km) on 4 September and between 175 and 140 hPa (12.9–14.4 km) on 5 September. The lighter shades of orange and blue show the original values and the darker shades the CO values corrected for oxidation by OH. a is the fitted slope and ci is the 95 % confidence interval of the fitted slope.

375   3.4   CO stable isotope composition

The LISA sampler CO mole fractions and stable isotope analysis results are presented in Table 4. Two types of samples can be distinguished, a plume sample and a background sample. It can been seen from Table 4 that the difference in their mole fraction and δ^{18}O is ~~most pronounced~~pronounced, whereas the δ^{13}C values are similar.

Table 4. LISA observations of CO~~and~~, mole fractions and the carbon and oxygen isotopic composition of CO, of the plume, P, and background, B. Here δ^{13}C(CO) is reported in ‰ ~~with~~ vs VPDB ~~as a reference material~~ and δ^{18}O(CO) in ‰ ~~with~~ vs VSMOW~~as a reference material~~.

	Altitude (km)	θ (K)	CO (ppb)	~~()~~ δ^{13}C(CO) ‰	δ^{18}O(CO) ‰
P (05 Sep)	13.6	370.3	74	~~405.5~~ -28.8	4.3
B (06 Sep)	13.4	368.9	34	~~405.2~~ -29.6	-1.0

The sample taken on 6 September can be considered as a background value for two reasons. First of all, the mole fraction
380   measurements from both LISA and AirCore agree with those of normal NH CO stratospheric mole fractions (Hoor et al., 2005). Secondly, the δ^{13}CO agrees to within ± 1 ‰ compared to the measurements performed in the southern hemisphere lowermost

stratosphere (Brenninkmeijer et al., 1995). Note that tropospheric CO and its isotopic composition exhibit a latitudinal gradient and a seasonal cycle, related to the OH-sink. It is not exactly known whether and to what extent these gradients exist in the stratosphere, an thus the agreement may be incidental.

385 The δ^{18}CO in the southern hemisphere SH is about 7.2 ‰ more depleted compared to the background value found from the observation presented here. Brenninkmeijer et al. (1995) attributed the relatively low values of ^{18}O to two determining factors; the unknown but probably low source signature for oxygen of methane-derived CO and the inverse kinetic isotope effect in the reaction with OH that depletes CO in ^{18}O. Nonetheless, lower δ^{18}O values are typically a sign of absence of any nearby sources other than oxidation of atmospheric methane (Brenninkmeijer et al., 1999).

390 The LISA CO mole fraction measured on 5 September compares best to the AirCore measurement of the day before and is in the middle of the plume; see Figure~~1. The~~ 1. We thus assume that the isotopic composition of that particular sample is representative for that in the plume. A comparison of the different δ^{18}O values, ~~see Table4 clearly show~~ in Table 4, clearly shows that the CO plume is ~~of different origin than stratospheric~~different than the background. Though the observed difference in δ^{13}C between the plume and the background sample is significant (0.8 ‰), ~~it cannot be excluded that this is due to~~ but could also be the result of natural variability. Furthermore, many sources carry a comparable δ^{13}C signature, see Table 1.

3.5 Source signature based on isotopic composition of CO

~~The observed enhancement carries a different isotopic composition compared to background air (Table 4). The additional information gained by isotopic analysis is useful since sources tend to produce products with a very distinct isotopic composition, which acts as a source signature.~~

400 Determining the source signature of the enhancement is not straightforward because of two reasons. First, the plume sample was affected by mixing in the troposphere during up-draft, and by mixing in the stratosphere. Secondly, the plume ~~is too old to ignore oxidative loss of carbon monoxide~~CO has been altered significantly by reaction with OH over the 25 day transport time. Thus, as explained in Section 2.3.3 the simple Keeling approach does not apply here.

Both the tropospheric background CO in NH summer (Bergamaschi et al., 2001; Mak et al., 2003) and stratospheric back-
405 ground CO, see Table 4, are more depleted in ^{18}O than the ~~sample obtained~~ CO in the plume. Thus, mixing would decrease the δ^{18}O of the plume. Since tropospheric and stratospheric δ^{13}C(CO) ~~were~~ are alike (see Table 3), δ^{13}C is not obviously affected by mixing.

In addition to mixing in the stratosphere, the ~~time-scale of transport is long enough for significant removal by~~ reaction with OH ~~to be important. The isotopic composition of the plume is subjected to~~ is associated with an inverse isotope effect at
410 stratospheric pressure for both ^{13}C and ^{18}O, depleting the remaining CO in both ^{13}C and ^{18}O (Röckmann et al., 1998). Thus, both removal by OH and mixing make the plume CO more depleted in ^{18}O ~~;~~ whereas the observed δ^{18}O value in the plume is higher than in the background. The δ^{13}C value is mainly affected by OH. The plume isotopic composition was thus originally more enriched in both ^{18}O and ^{13}C.

Equation 4 is used to estimate the CO mole fraction and oxygen and carbon isotopic composition of the plume 25 days before
415 the observations as a function of the unknown stratospheric fraction of air mixed into the sample, f_{strat}. It is assumed that pres-

sure and temperature remained constant during transport. The number density of OH assumed for the stratosphere is ~~$1.2 \cdot 10^6$~~ ~~and is as used to derive an estimate of the enhancement ratio in Section 3.3. The estimated mole fraction and the carbon and oxygen isotopic composition are shown in Figure 5 for different number densities. The~~ $n(\mathrm{OH}) = 1.2 \cdot 10^6$ cm^{-3}. The model re-

sults show that the isotopic composition of the plume would have been ~~$-27.5\,‰ \le \delta^{13}\mathrm{C} \le -25\,‰$ and $10\,‰ \le \delta^{18}\mathrm{O} \le 16\,‰$~~ $-27.5\,‰ \le \delta^{13}\mathrm{C} \le -26.6\,‰$ and $10\,‰ \le \delta^{18}\mathrm{O} \le 14\,‰$ after injection into the stratosphere, see Figure~~5~~ 5. This range corresponds to the modelled results for $n(\mathrm{OH}) = 1.2 \cdot 10^6$ cm^{-3}. Figure 5 also shows results of the same calculations assuming twice higher OH number densities and no oxidation by OH.

Using the OH-corrected estimate of ~~$f_{\mathrm{strat}} = 0.66$~~ $f_{\mathrm{strat}} = 0.55$ from the Monte-Carlo simulation in Section 3.6, the plume $\delta^{13}\mathrm{C}$ and $\delta^{18}\mathrm{O}$, directly after injection in the stratosphere, are estimated to be ~~-25.6~~ $-27.0\,‰$ and ~~11.7~~ $11.4\,‰$ respectively. Hence, the fractionation that occurred during transport depleted $^{13}\mathrm{C}$ by ~~3.2~~1.8 ‰ and ~~7.4~~7.1 ‰ $^{18}\mathrm{O}$.
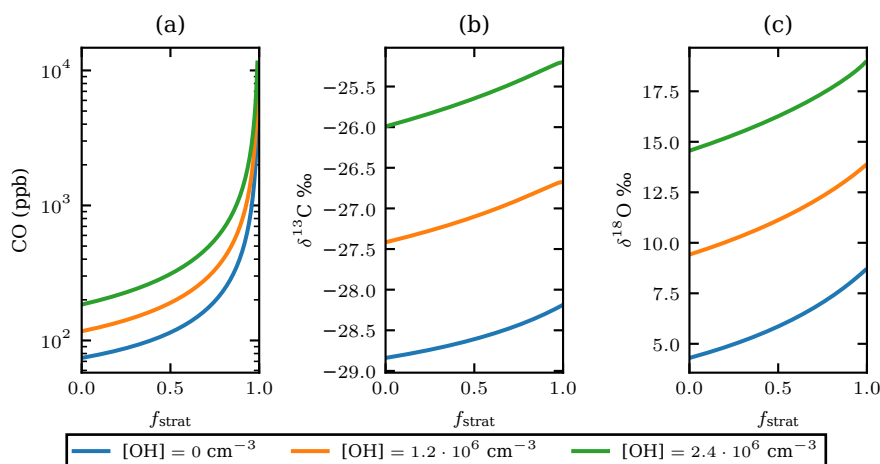

Figure 5. The estimated CO mole fractions on a logarithmic scale (a), ~~note~~ the ~~logarithmic scale on the y axis, and the~~ carbon (b), and oxygen (c) isotopic composition of the plume CO for different OH number densities versus f_{strat} just after injection into the stratosphere. OH number densities are shown in the legend, the value of $1.2 \cdot 10^6$ can be considered representative of the stratosphere. The value of $2.4 \cdot 10^6$ is arbitrarily added, as twice the stratospheric value and serves as an upper estimate.

Note that when the number density of OH is 0 cm^{-3} Equation 4 describes simple end-member mixing with rate k_{er}. If the Keeling method is applied, a source signature of $\delta^{13}\mathrm{C}_{\mathrm{src}} = -28.1$ and $\delta^{18}\mathrm{O}_{\mathrm{src}} = 8.7$ is obtained, ~~suggesting~~ providing a lower limit of the source signature. ~~It can be seen from Figure 5 that even with~~ Upper limits for the source signature are derived assuming a very large amount of stratospheric mixing and high OH number density the source signature has $\delta^{13}\mathrm{C} < -22\,‰$ and $\delta^{18}\mathrm{O} < 21\,‰$. Based on this analysis the source signature of the plume origin has $-27.5\,‰ \le \delta^{13}\mathrm{C} \le -22\,‰$ and $10\,‰ \le \delta^{18}\mathrm{O} \le 21\,‰$.

The analysis above shows that the plume was initially more enriched in both $^{13}\mathrm{C}$ and $^{18}\mathrm{O}$ than at the time of observation. Comparing ~~this~~ the derived signature to the source signatures in Table 1, it is clear that the source signature is similar to that of

CO produced in wildfires. In fact, qualitatively, the plume was more enriched, then computed in Figure 5, if we would correct for the tropospheric air mixed into the air parcel. Then the values for the troposphere presented in Table 3 have to be assumed, and the source signature comes close to the expected range of isotopic compositions for the wildfire presented in this study, Table 3.

Fossil fuel combustion sources, the only other source that produces CO containing higher $\delta^{18}O$, can be excluded for two reasons. First, the source signature of high temperature combustion, $\sim 23.5\,‰$, ~~would require an unusually high fractionation to explain the observed value~~ is outside the calculated range of the source signature. Secondly, the enhancement ratio of $\Delta CO : \Delta CO_2$ is too high for modern day fossil fuel combustion ~~.~~ (see e.g. Popa et al. (2014)).

On the basis of the observed $\delta^{13}C$ signature, CH_4-oxidation can be excluded as a source, as methane-derived $\delta^{13}C$ is usually far more depleted (Brenninkmeijer et al., 1999). Finally, oxidation of NMHCs can be excluded as a significant source. The total amount of NMHCs in the stratosphere is on the order of several ppt (Scheeren et al., 2003). Furthermore, estimates of the NMHCs source signature suggests that the oxygen signature is in the range 0–3.6 ‰ (Brenninkmeijer and Röckmann, 1997; Vimont et al., 2019). The NMHCs produced in the fire have a mean enhancement ratio to CO of the order of $ppt\,ppm^{-1}$ (Mauzerall et al., 1998), and thus would result in a very small in situ source of CO that ~~has~~ would have a very small effect on the isotopic composition.

3.6 ~~Stratosphere-troposphere exchange estimate~~ Estimate of the tropospheric air fraction based on the in situ observations

The fractions of tropospheric, f_p, and stratospheric air, f_s, in the plume were determined using Equation ~~13 was used to determine the fractions of tropospheric and stratospheric air in the plume~~ 13. The results of two Monte-Carlo simulations are shown in Figure 6, one simulation with CO observation corrected for OH, and one without the correction. The ~~mode~~ mean of each distribution suggest that the tropospheric air fraction is ~~46~~48 $\pm$ 21 % and the stratospheric fraction is ~~54~~52 $\pm$ 21 %. After the correction for oxidation, this shifts the tropospheric contribution to ~~34~~45 $\pm$ 21 %, and the stratospheric contribution to ~~66~~55 $\pm$ 21 %. Thus, ignoring the oxidation results in a small bias in the estimated stratospheric and tropospheric contribution to the plume.

4 Discussion

~~In this study, it is shown that stable isotope analysis and mole fractions obtained by LISA and AirCore indicate the presence of wildfire smoke in the stratosphere. Using the stable isotope and mole fraction observation, the contribution of tropospheric air and stratospheric air to the composition of the plume is estimated.~~

4.1 Enhancement ratios and plume age

Initially, the plume was observed from a clear CO mole fraction increase in the stratosphere, present in two AirCore profiles. The CO plume mole fraction measurements, up to 90 ppb, are lower compared to other plumes measured in the stratosphere,
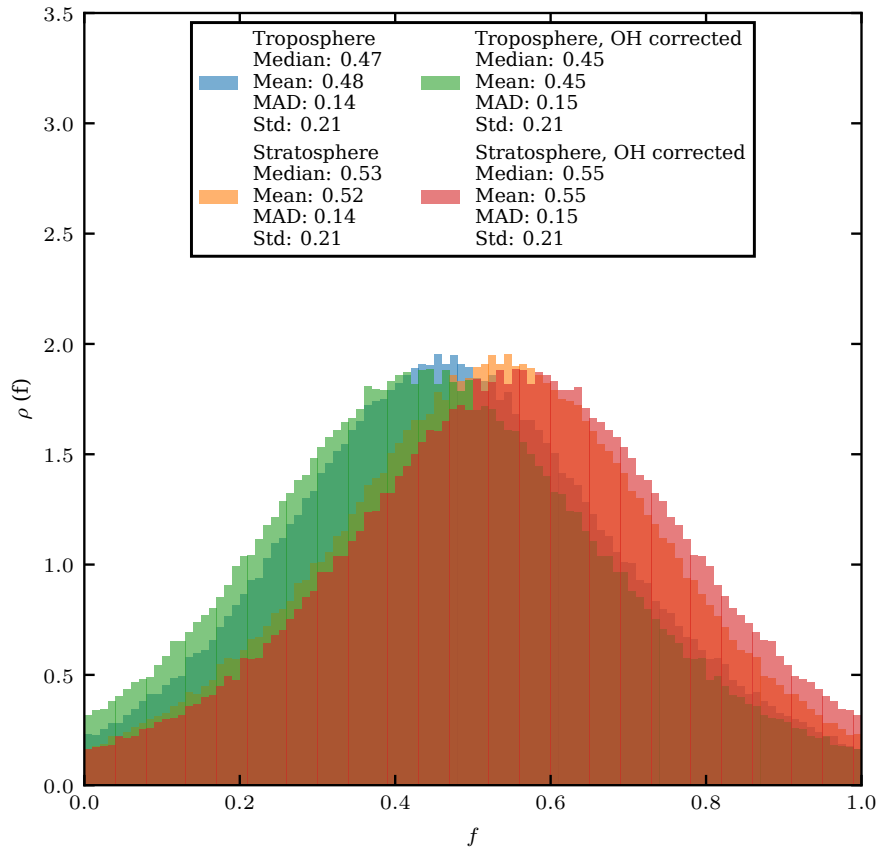
Figure 6. Probability density of the ~~airmass~~ air fraction by volume, as a result of a Monte-Carlo simulation. ~~Randomly selected input parameters, the end-member definitions presented in Table 3 were used to define Equation 13, which was consequently solved for the airmass fractions by volume f. This process was continued until $1 \cdot 10^6$ solutions were obtained, i.e. $1 \cdot 10^6$ solutions passed the filter (see Table 2.3.4).~~ The data is binned into 1 % bins. ~~A second run was performed~~ Results with and without a correction for ~~the oxidation applied to the observation, AP~~ OH ~~-corrected in Table 3~~ are shown. ~~The mode of each distribution was obtained from a fit to the distribution.~~

notably by Waibel et al. (1999) (300 ppb), Jost et al. (2004) (200 ppb), and by Cammas et al. (2009) (250 ppb). First, the plume reported is older than other observations. The estimated plume age was 25 days, where the other observations were sampled after 7 to 14 days. Hence, the plume observed here was affected more by mixing and photo-chemistry in the stratosphere. Secondly, it is ~~not possible to determine from the data presented here, whether~~ unlikely that the accidental encounter sampled the centre of the plume ~~was sampled, in the horizontal sense,~~ where CO mole fractions are highest. Finally, it should be pointed out that the plume was encountered during AirCore vertical profiling which has limited vertical resolution ~~. The vertical resolution of the AirCore measurements was estimated to be about 150 between 13 and 14~~ which could smooth out the maximum values.

In addition to the anomalous CO mole fraction, enhancement ratios of CO vs CO_2 obtained from the regression analyses, based on concurrent AirCore CO_2 and CO measurements, agree with the ratios ~~measured previously in~~ obtained in previous studies on biomass burning emissions. The observed ratios of 34–40 $\mathrm{ppb\,ppm}^{-1}$ in this study is lower than the measured enhancement ratios of 50 $\mathrm{ppb\,ppm}^{-1}$ by Jost et al. (2004) and 48–73 $\mathrm{ppb\,ppm}^{-1}$ by Andreae et al. (2001). The plumes reported by both Jost et al. (2004) and this study originated from forest fires in North America (40°–55° N), and were observed at a similar altitude of approximately 1.5 km above local tropopause. It is a reasonable assumption that the initial enhancement ratios and the mixing of the plumes with lower stratospheric backgrounds are similar. However, the plume age of roughly 25 days in this study is significantly older than that of 10–14 days observed by Jost et al. (2004). Therefore, the ageing of the plume coupled with the OH-related destruction of CO likely explains the difference in the observed ratios. Similarly, the age of 9–10 days of the plumes observed by Andreae et al. (2001) is also significantly younger than the plume age found in this study. ~~Furthermore, their observation was made in tropical tropospheric air, and the background mole fraction of about 100 is significantly higher than the stratospheric background of roughly 40 ppb, which contributes to their higher observed enhancement ratios than those in this study.~~ Indeed, the OH-corrected enhancement ratios of 53–62 $\mathrm{ppb\,ppm}^{-1}$ in this study come closer to the similar OH-corrected enhancement ratios of 64–98 $\mathrm{ppb\,ppm}^{-1}$ in Andreae et al. (2001), and the remaining difference may be caused by the different type of fuels of biomass burning, forest in this study and savannah/forest in Andreae et al. (2001).

~~This ratio is comparable to the measured high-altitude enhancement ratio, 50 and 55 Jost et al. (2004) however lower than those observed by Andreae et al. (2001).They found 74 for a higher altitude plume after correction for plume ageing.There are two important differences between the measurements in this study and the measurements presented by Andreae et al. (2001). First, their observation was made in tropical tropospheric air , which is affected by different photo-chemistry and transport properties than the stratospheric measurements presented here. Secondly, they found a lower age of 9.5 for the observed plume which makes the correction for removal by OH less uncertain~~The enhancement ratios computed directly are much higher than the results discussed above. First, it is difficult to assume what background should be used. The stratospheric background mole fractions, especially the CO_2, varies with altitude and in time, e.g. comparing the AirCore profiles from 6 and 7 September, which questions the assumption of a constant background. It is thus difficult to state with certainty that data from 6 September is representative for the background. However, as the plume, and air directly adjacent to it, move together, it can be assumed that the air surrounding the plume has constant mole fractions. The best estimate of the mole fractions are those measured

above and below the plume, obtained from the vertical profile. Hence, the most reliable estimate of the enhancement ratio is obtained from the regression analysis discussed above. On a final note, the very small standard deviation of 1–2 ppb obtained for 4 September is largely due to the small amount of data, $N = 5$, and the large averaging window of 25 data points.

4.2 Isotopic composition of plume-CO

The stable isotope source ~~signature~~ signatures of CO ~~qualitatively~~ supports wildfire smoke as the source. ~~Very little stable~~
505 ~~isotope measurements exist on stratospheric air. On the other hand, several tropospheric measurements of other wildfire events were published so a comparison can be made with those.~~

~~First, the~~ The reported δ^{18}O source signature for Siberian boreal forest fires, 14.8 ‰, and 9.0 ‰, respectively (Bergamaschi et al., 1998; Tarasova et al., 2007), compare well with the observation made in this work after estimating the fractionation that occurs during oxidation, see Section~~3.5. Although the transport-history of the observations presented here complicates the~~
510 ~~determination of the source signature, it was shown that the original δ^{18}O was higher.~~ 3.5.

~~In addition, Brenninkmeijer and Röckmann (1997) found~~ Brenninkmeijer and Röckmann (1997) reported a source signature of 4.5 ‰ for wildfire smoke, based on southern hemisphere observations. As argued by Brenninkmeijer and Röckmann (1997) the samples must have been affected by the strong fractionation accompanying the reaction of CO with OH along the mixing, something that was not accounted for in their method used to derive the source signature. The lifetime of CO against removal
515 by OH in the stratosphere is considerably longer than in the troposphere, and hence the wildfire sample presented here was likely less affected by fractionation than their measurement.

It is shown that CO stable isotope measurements can help pollution events in the stratosphere to be identified. It must be noted that this study took advantage of the fact that, on the days following the pollution event, clean background air was sampled. Thus, a direct comparison of background air and ~~pollution~~ polluted air was possible. Without the measurement of
520 background air, source attribution would have been difficult from stable isotope measurements, as little is known about the CO isotopic composition. Fundamental knowledge of CO isotopic composition and its temporal and latitudinal variation in the stratosphere is vital for the detection of future pollution events based on CO measurements.

In addition to a poorly understood isotope budget of the stratosphere, ~~studies would benefit from measurements of source signatures from lab experiments. Although methane-derived can be discriminated based on , most of the other sources are~~
525 ~~similar in . The~~ the case made in this study would have been stronger, if the oxygen signature of both methane and NMHCs were known more precisely. Our fundamental knowledge of the CO isotopic composition in the stratosphere would also benefit from those measurements, as methane is the main source of CO in the stratosphere.

4.3 Assessment of tropospheric and stratospheric airmass contributions

Finally, the ~~airmass-fraction~~ airmass-fractions of the troposphere and stratosphere were derived using the tracer observations.
530 The results suggest that the 2017 pyro-Cb plume observed above Sodankylä consist of approximately ~~34 ± 14~~45 ± 21 % tropospheric air polluted with wildfire smoke. This is ~~in qualitative agreement with~~ comparable to the model simulations from Trentmann et al. (2006) on the Chisholm fire in 2001, an event similar to the 2017 British Colombia fires. ~~Yet another event~~

~~was modelled by Cammas et al. (2009) and~~ Cammas et al. (2009) modelled their observations of wildfire smoke from several fires in Canada and Alaska and they estimated the amount of polluted boundary layer air above the tropopause to be 15–20 %.

5 Conclusions

A wildfire smoke plume in the lower stratosphere ~~is~~ was investigated using in situ observations CO and CO_2 from AirCore, and stratospheric $\delta^{13}C$ and $\delta^{18}O$ in CO from LISA. The plume was identified by enhanced CO mole fractions at approximately 13.6 km altitude, present on two consecutive days, and extending over 1 km in altitude. The plume's enhancement ratio of CO to CO_2 mole fractions was in the range ~~34–40~~52–62 ppb ppm^{-1}. The stable isotopic composition of carbon and oxygen in CO support wildfire smoke as the source for the enhanced CO mole fractions observed in both AirCore and LISA samples. Using the CLaMS back-trajectory module and CALIOP backscatter data the source region is determined to be British Colombia, Canada. The smoke was injected on 12 August 2017, 24–25 days before the observations were made. The age of the plume aided in the estimation of the amount of oxidation, a 35–37 % loss of CO, and the accompanying isotopic fractionation, ~~3.3~~1.8 ‰ for $\delta^{13}C$ and ~~7.6~~7.1 ‰ $\delta^{18}O$. Using this information, the enhancement ratios corrected for oxidation ranged from 53 to 62 ppb ppb^{-1}. The plume isotopic composition of oxygen and carbon in CO was estimated to be ~~−25.6~~−27.0 ‰ and ~~11.7~~11.4 ‰. From the LISA observations, it was possible to determine the fractions of tropospheric, ~~34±14~~45±21 %, and stratospheric air, ~~66±14~~55±21 %, in the plume using a three end-member mixing model.

Appendix A: Derivation of Equation 6 and Equation 8

A constant entrainment rate is assumed, i.e. so that

$$\frac{dV}{dt} = kV \tag{A1}$$

This can be solved to yield:

$$V(t) = V_0 \exp(kt) \tag{A2}$$

with V_0 the initial volume of the air-parcel containing the contamination. Thus $V_0/V_t = f_{plume}$ and

$$V(t) = f_{plume}V_{plume} + f_{strat}V_{strat} \tag{A3}$$

with $f_{plume} + f_{strat} = 1$. Then Equation A2 can be written as follows:

$$1 - f_{strat} = \exp(-kt) \tag{A4}$$

this can be rearranged to give the end result, k in terms of f_{strat}:

$$k = \frac{-\ln(1 - f_{strat})}{t} \tag{A5}$$

Starting from

$$\frac{dn(CO)}{dt} = -k_{er}\left(n(CO) - n_{bg}(CO)\right) - k_1 n(OH) n(CO) \tag{A6}$$

Letting $n(CO) = x$ and:

$$a = -k_{er} - k_1 n(OH) \tag{A7}$$

and

565    $$b = k_{er} n_{bg}(CO) \tag{A8}$$

After substitution and rearranging would result in:

$$\frac{dx}{ax + b} = dt \tag{A9}$$

~~Integration yields~~

Integration yields:

570    $$\frac{1}{a}\ln(ax + b) + C = t \tag{A10}$$

where C is an integration constant. Solving for x gives:

$$x(t) = \frac{1}{a}\left(\exp(at)\exp(-Ca) - b\right) \tag{A11}$$

$\exp{-Ca} =$ constant so it can be replaced by yet another arbitrary constant c.

$$x(t) = \frac{1}{a}\left(c\exp(at) - b\right) \tag{A12}$$

575   After doing so, the constant c can be determined from the boundary condition $x(t = 0) = x_0$:

$$x_0 = \frac{c - b}{a} \tag{A13}$$

which is equivalent to:

$$c = ax_0 + b \tag{A14}$$

substituting Equation A14 into Equation A12 yields:

580    $$x(t) = \left(x_0 + \frac{b}{a}\right)\exp(at) - \frac{b}{a} \tag{A15}$$

~~Finally Equation~~Finally Equation A7 and Equation A8 and $n(CO) = x$ can be used to obtain:

$$n(CO)(t) = \left(n_0(CO) + \frac{k_{er}n_{bg}(CO)}{-k_{er} - k_1 n(OH)}\right)\exp\left(-\left(k_{er} + k_1 n(OH)\right)t\right) - \frac{k_{er}n_{bg}(CO)}{-k_{er} - k_1 n(OH)} \tag{A16}$$

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