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## Effect of mixing structure on the water uptake of mixtures of

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## ammonium sulfate and phthalic acid particles

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Abstract. Aerosol mixing state regulates the interactions between water molecules and particles and thus controls the aerosol activation and hygroscopic growth, which thereby influences the visibility degradation, cloud formation, and its radiative forcing. Current studies on the mixing structure effects on aerosol hygroscopicity, however, are few reported. Here, we investigated the hygroscopicity of AS/PA aerosol particles with different mass fractions of PA in the different mixing states in terms of initial particle generation. Firstly, effect of phthalic acid (PA) coatings on the hygroscopic behavior of the core-shell-generated mixtures of ammonium sulfate (AS) with PA has

24	been studied using a coating-hygroscopicity tandem differential mobility analyzer (coating-
25	HTDMA). The slow increase in the hygroscopic growth factor of core-shell-generated particles is
26	observed with increasing thickness of coating PA prior to DRH of AS. At RH above 80 %, a decrease
27	in hygroscopic growth factor of particles occurs as the thickness of PA shell increases, which
28	indicates that the increase of PA mass fractions leads to a reduction of the overall core-shell-
29	generated particle hygroscopicity. In addition, the use of the ZSR relation leads to the
30	underestimation of the measured growth factors of core-shell-generated particles without
31	consideration of the morphological effect of core-shell-generated particles, especially at higher RH.
32	Secondly, in the case of the AS/PA initially well-mixed particles, a shift of deliquescence relative
33	humidity (DRH) of AS (~80 %, Tang and Munkelwitz (1994)) to lower relative humidity (RH) is
34	observed due to the presence of PA in the initially well-mixed particles. The predicted hygroscopic
35	growth factor using the ZSR relation is consistent with the measured hygroscopic growth factor of
36	the initially well-mixed particles. Moreover, we compared and discussed the influence of mixing
37	states on the water uptake of AS/PA aerosol particles. It is found that the hygroscopic growth factor
38	of the core-shell-generated particles is slightly higher than that of the initially well-mixed particles
39	with the same mass fractions of PA at RH above 80 %. The observation of AS/PA particles may
40	contribute to a growing field of knowledge regarding the influence of coating properties and mixing
41	structure on water uptake.

# **1 Introduction**

44 The ability of aerosol particles to absorb and maintain water molecules, called hygroscopicity, is45 one of the most important physicochemical properties of atmospheric aerosol particles with

46	profound implications (Shi et al., 2012; Lei et al., 2014, 2018; Gupta et al., 2015; Hodas et al., et
47	al., 2015; Zawadowicz et al., 2015; Martin et al., 2017). It might determine the phase state (Mu et
48	al., 2018), size (Peng et al., 2001; Choi et al., 2002), mixing state, optical properties, and chemical
49	reactivity of atmospheric aerosols exposed to the environment of the different RHs (Heintzenberg
50	et al., 2001; Rudich et al., 2003; Spindler et al., 2007; Abo Riziq et al., 2007, 2008; Eichler et al.,
51	2008). Moreover, the change of these properties after water absorption on aerosol particles can
52	strongly affect the cloud formation, aerosol radiative forcing, global climate, and even human health
53	(Cheng et al., 2008; Reutter et al., 2009; Rose et al., 2011; Stock et al., 2011; Liu et al., 2012, 2013;
54	Tie et al., 2017). Therefore, the interaction between water molecules and aerosol particles is crucial
55	for a better understanding of the aerosol-cloud-climate effects in the atmosphere (Sjogren et al.,
56	2007; Zamora et al., 2011; Jing et al., 2016).
57	Atmospheric aerosols contain a complex mixture of inorganic and organic compounds in the
58	different mixing structures, e.g., externally mixed, internally mixed (Ganguly et al., 2006). The
59	
	internally well-mixed aerosol particles may be divided into homogeneous and heterogeneous
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60 61 62 63	internally well-mixed aerosol particles may be divided into homogeneous and heterogeneous internally mixed aerosol particles (Lang-Yona et al., 2009), which could, in turn, strongly influence the water uptake, optical properties, and the cloud condensation nuclei (CCN) ability of the particles (Lesins et al., 2002; Falkovich et al., 2004; Zhang et al., 2005; Schwarz et al., 2006; Su et al., 2010). Most of the previous studies on the hygroscopic behavior of multi-components aerosol focus on the
<ul> <li>60</li> <li>61</li> <li>62</li> <li>63</li> <li>64</li> </ul>	internally well-mixed aerosol particles may be divided into homogeneous and heterogeneous internally mixed aerosol particles (Lang-Yona et al., 2009), which could, in turn, strongly influence the water uptake, optical properties, and the cloud condensation nuclei (CCN) ability of the particles (Lesins et al., 2002; Falkovich et al., 2004; Zhang et al., 2005; Schwarz et al., 2006; Su et al., 2010). Most of the previous studies on the hygroscopic behavior of multi-components aerosol focus on the well-mixed particles generated from internally mixed solutions (Miñambres et al., 2010; Shi et al.,

- 66 (2002) studied on the effects of glycerol, succinic acid, malonic acid, citric acid, and glutaric acid
- 67 on the hygroscopic properties of sodium chloride and AS in the initially well-mixed aerosol particles,

68	respectively, using an electrodynamic balance. They observed that the deliquescence and
69	efflorescence of sodium chloride and AS were affected by the presence of different organic
70	components in the mixed aerosol particles. Concerning the hygroscopicity of the heterogeneity of
71	internally mixed aerosol particles, such as core-shell particles, there are several studies on
72	investigating their interaction with water molecules (Ciobanu et al., 2009; Song et al., 2012;
73	Shiraiwa et al., 2013; Hodas et al., 2015; Song et al., 2018). However, to date, few laboratory studies
74	have been investigated on the influence of organic coatings on the hygroscopic behavior of core-
75	shell particles and the difference of mixing state effects on the hygroscopicity of aerosol particles
76	(Zhang et al., 2008; Pagels et al., 2009; Xue et al., 2009; Lang-Yona et al., 2010; Ditas et al., 2018).
77	E.g., hygroscopicity tandem differential mobility analyzer (HTDMA) study on the organic coating
78	effects on the hygroscopicity of core-shell aerosol particles was studied by Maskey et al. (2014).
79	They observed a shift of DRH of AS to lower RH for the core-shell particles due to presence of
80	levoglucosan coatings. They further compared water absorption on AS/levoglucosan core-shell
81	particles and the AS/succinic acid core-shell particles. They suggest that different organic coatings
82	lead to changes in the hygroscopic properties of core-shell particles. Chan et al. (2006) investigated
83	hygroscopicity of AS coated with different mass fractions of glutaric acid during two continuous
84	humidification and dehumidification cycles using a Raman spectra and an electrodynamic balance.
85	They observed different hygroscopic behavior and morphology of aerosol particles between the two
86	humidification and dehumidification cycles due to the different mixing states. Therefore, to
87	investigate organic coating effect on the hygroscopicity of aerosol particles and further to study on
88	difference of mixing states effects on the hygroscopic behavior of aerosol particles are crucial for
89	estimation the direct and indirect radiative effect of aerosol particles on the Earth's climate (Saxena

90 et al., 1995; Ansari et al., 2000; Maskey et al., 2014).

91	PA is ubiquitous in rural mountains and marine atmosphere in Asia (Wang et al., 2011). It is mainly
92	produced by the photo-oxidation of volatile organic compounds (VOCs), such as xylene, and
93	naphthalene (Kawamura and Ikushima, 1993; Schauer et al., 1996; Zhang et al., 2016). Also, PA has
94	been identified as a significant contributor to the urban organic compounds (Rogge et al., 1993). PA
95	particles are generally used as a tracer for the secondary organic aerosol (SOA) in atmospheric fine
96	particles (Schauer et al., 2000, 2002). Furthermore, these aromatic acids like PA, which are photo-
97	oxidation products of anthropogenic precursors such as toluene, xylene, and naphthalene, are
98	initially formed as gaseous products and subsequently condensed onto pre-existing particles. Also,
99	Kleindienst et al. (1999) investigated that hygroscopic particles formed after irradiating toluene, p-
100	xylene, and 1,3,5-trimethylbenzene in the presence of NOx and AS seed in a chamber. This suggests
101	the likely existence of PA in atmospheric aerosol particles. Organic coatings on inorganic aerosol
102	particles in the atmosphere can play an important role in the range of RHs over which particle-bound
103	water influences aerosol properties, such as the overall density, the light scattering behavior and the
104	refractive index. Recently, Zhang et al. (2016) reported the importance of atmospheric PA aerosol
105	particles in the visibility degradation and the formation of CCN. The organic PA can have a profound
106	effect on light scattering, hygroscopictiy, and phase transition properties of multicomponent
107	atmospheric aerosols. However, these physicochemical properties of PA have been little
108	documented in the literature (Brooks et al., 2004; Liu et al., 2016). Here, we summarized a few
109	studies on the hygroscopicity of the PA-containing aerosol particles. Brooks et al. (2004)
110	investigated continuous hygroscopic growth of PA aerosol particles in the humidification process
111	using HTDMA technique. Hori et al. (2003) and Huff Hartz et al. (2006) measured the high CCN

112	activity of PA in spite of its low solubility. Also, liquid-liquid phase separations (LLPS) with aerosol
113	particles consisting of organic and inorganic components was observed by many groups (Ciobanu
114	et al., 2009; Betram et al., 2011; Song et al. 2012a, 2012b; You et al., 2014). For example, Song et
115	al. (2012a, 2012b) investigated that LLPS occurs in the mixed dicarboxylic acids containing 5, 6,
116	and 7 carbon using an optical microscopy and micro-Raman spectroscopy, and further established
117	that occurrence of LLPS of aerosol particles has an average elemental oxygen-to-carbon (O:C) ratio
118	of the organic fraction of less than 0.8. Subsequently, the occurrence of liquid-liquid phase
119	separation in the internally mixed aerosols consisting of AS and PA was performed by Zhou et al.
120	(2014) during the dehumidification processes using a Raman spectra. You et al. (2014) further found
121	that the LLPS of aerosol particles with a different O: C ratio of 0.5 <o: c<0.8="" depends="" on="" td="" the="" types<=""></o:>
122	of organic functional groups and inorganic salts presented. Therefore, these studies suggest that the
123	LLPS in the mixed organic and inorganic components aerosol particles is influenced by the amounts
124	of organic and inorganic aerosol components and types.
125	In this work, AS is chosen as a test substance because it is a major component of the atmospheric
126	aerosol and its thermodynamic behavior is well characterized. Most importantly, ammonium sulfate
127	particles are stable and not volatile, useful features for HTDMA studies. We investigated the effect
128	of the different thickness of coating PA and the AS core size on the water uptake of core-shell-
129	generated particles, and further studied the effect of the mixing states on the hygroscopic behavior
130	of PA/AS aerosol particles using HTDMA technique. For example, we compared the hygroscopic
131	behavior of initially well-mixed AS/PA particles with that of core-shell-generated AS/PA particles
132	with the same PA mass fractions. In addition, we used the Zdanovskii-Stokes-Robinson (ZSR)
133	relation to predict the hygroscopic growth factor (GF) of mixed aerosol particles in the different

mixing structure. Moreover, the Aerosol Inorganic-Organic Mixtures Functional groups Activity
Coefficients (AIOMFAC) model (Zuend et al., 2008; 2011) with a version of the liquid-liquid
equilibrium (LLE) algorithm was employed in our study to predict the phase compositions of liquid
and solid phases for a given composition of a mixture (Zuend and Seinfeld. 2013).

138

#### 139 **2** Experimental and modeling methods

#### 140 **2.1 HTDMA setup and experimental protocol**

141 HTDMA setup is employed to measure the aerosol nanoparticle hygroscopic growth factor ( $g_f$ ) and phase transition in the RH range from 5 % to 90 %. Here,  $g_f$  is defined as the ratio of mobility 142 143 diameter of aerosol particles after humidification ( $D_m(RH)$ ) to that dry condition ( $D_m(<5 \% RH)$ ). 144 Figure 1 shows a schematic diagram of the HTDMA setup. It is comprised of four main components, 145 including three differential mobility analyzers (DMA), a condensation particle counter (CPC), a 146 humidification system, and a coating system. The more detailed information on the HTDMA setup, 147 calibration, and verification have been described elsewhere (Lei et al., 2014; 2018; Jing et al., 2016; 148 Liu et al., 2016). In our study, the particle-sizing, the aerosol/sheath flow rates, and DMA voltage 149 supply have been calibrated every month, respectively. The uncertainty of aerosol/sheath flow rates 150 are kept within  $\pm 1\%$  around the reference values. The deviations of the measured DMAs voltage 151 from set-point values is less than  $\pm 1\%$ . The sizing agreement of DMAs between the measured diameter of polystyrene latex (PSL) spheres and their nominal diameter ( $100\pm3$  nm) is within  $\pm1\%$ . 152 153 Thus, the calculated uncertainty of growth factor depends on the error propagation formula by

$$\sqrt{\left(\left(g_f \frac{\sqrt{2}\varepsilon_{Dp}}{D_p}\right)^2 + \left(\varepsilon_{RH} \frac{dg_f}{dRH}\right)^2,\right)}$$
(Mochida and Kawamura, 2004). Here,  $\varepsilon_{Dp}$ ,  $\varepsilon_{RH}$ , and  $g_f$  are

uncertainty of particle mobility diameter, relative humidity, and growth factor at different RH, respectively. The average sizing offsets of our system is taken here as  $\frac{\varepsilon_{Dp}}{D_p}$ . Also, the RH uncertainty is the accuracy of RH sensor (±2 %). In this study, the calculated uncertainty of growth factor of AS/PA aerosol particles is ~1%-2%. The chemical substances and related to physical properties are available in the Supporting Information Table S1. The solutions used in our measurements are 160 prepared with distilled and de-ionized Milli-Q water (resistivity of 18.2 MΩ cm at 298 K).

#### 161 2.1.1 Initially well-mixed AS/PA Aerosol particles

162 Briefly, poly-disperse aerosol particles were atomized from homogeneous bulk solutions with 163 different mass fractions of PA and AS (Fig. 1), assuming that the compositions of aerosol particles 164 remain the same as that of the solutions used in the atomizer (MSP 1500, MSP). Due to morphology 165 and mixing state of AS/PA aerosol particles generated by an initially well-mixed aqueous solution as indicated by Fig. S2 at dry RH, note that aerosol particles generated this way are referred to as 166 167 initially well-mixed aerosol particles. The resulting particles were dried and subsequently charged 168 through a dryer and then a neutralizer, respectively. A mono-disperse distribution of particles with 169 a desired diameter were selected by the first differential mobility analyzer (DMA1) with RH below 5 %. After particle sizing, the aerosol particles were exposed to a humidification mode (5  $\% \rightarrow 90\%$ ) 170 171 in the Nafion conditioner tubes. The number size distributions of humidified aerosol particles were 172 then measured by a DMA3 coupled with a CPC. To have a precise control of the aerosol RH, the 173 flow rates of the humid and dry air were adjusted with a proportional-integral-derivative (PID) 174 system. Also, to ensure the sufficient water equilibrium with aerosol particles, the difference 175 between RH2 and RH3 (RH in the sheath flow) was within 2 % during the experiment.

#### 176 2.1.2 Core-shell-generated AS/PA aerosol aerosols

The AS core aerosol particles were generated from an aqueous solution of AS (0.05 wt %) by an atomizer. After a passage through a silica gel diffusion dryer and a neutralizer, the AS core aerosol particles with a certain diameter (100, 150, and 200 nm, respectively) were firstly selected by a DMA1 and then exposed to organic vapors in a coating system. To be specific, the coating system contains a controlled silicone oil bath vaporizer, a reservoir of organic compound, and a condenser.

182	The AS core particles passed through a sealed flask immersed in a silicone oil bath. The sealed flask
183	was filled with the PA powder. The PA vapors were enriched into the aerosol flow by heating. The
184	temperature required for vaporizing PA is between ~100 and ~130°C, which corresponds to coating
185	thickness between 10 and 50 nm. The resulting organic vapors were condensed onto the AS core
186	particles after cooling to an ambient environment through a condenser. Similarly, this system for
187	coating organic components on the particles has been proved to be efficient by Abo Riziq et al.
188	(2008). The coated particles of a certain size were then selected by the DMA2 to determine the
189	thickness of organic components ( $D_{total} = D_{core} + coating$ ). After core-shell-generated particle-sizing,
190	aerosols were pre-humidified in a Nafion tube and flowed into the second Nafion humidifier at the
191	set RH2 to reach equilibrium for growth of aerosol particles. Finally, the humidified core-shell-
192	generated particles were detected by a DMA3 and a CPC at ambient temperature. The uncertainty
193	of thickness of coating PA is $\pm$ 1.0 nm, considering the fluctuation in temperature and uncertainty
194	of sizing measurements by DMAs. Due to morphology and mixing state of AS/PA aerosol particles
195	generated by a coating-HTDMA at dry RH, note that aerosol particles generated this way are
196	referred to as core-shell-generated aerosol particles.

- 197 **2.2 Theory and modeling methods**
- 198 2.2.1 GF data fit
- 199 We use the following expression to predict the hygroscopic growth factor of individual components.

200 
$$GF = \left[1 + (a + b * a_w + c * a_w^2) \frac{a_w}{1 - a_w}\right]^{\frac{1}{3}}$$
 (1)

- Here it is assumed that water activity  $(a_w)$  is equal to the water saturation ratio  $(a_w = RH / 100 \%)$ .
- 202 The coefficients a, b, and c are determined by fitting Eq. (1), and their values are shown in Table 1
- 203 according to the measured GF data against RH. The equation (1) is expected to fit the continuous

water uptake behavior of particles (Brooks et al., 2004; Kreidenweis et al., 2015).

#### 205 2.2.2 GF predictions by ZSR

We use the Zdanovskii-Stokes-Robinson (ZSR) relation to calculate the hygroscopic growth factor of mixed particles,  $GF_{mixed}$ . The GF of a mixture can be estimated from the  $GF_j$  of the pure components *j* and their respective volume fractions,  $\mathcal{E}_j$ , in the mixture (Malm and Kreidenweis, 1997).  $R_{As}$  is radius of core and  $R_{core-shell}$  is radius of core-shell-generated aerosol.

210 
$$GF_{mix} = \left[\sum_{j} \varepsilon_{j} (GF_{j})^{3}\right]^{\overline{3}}$$
(2)

211 
$$\varepsilon_{AS} = \frac{\frac{4}{3}\pi R_{AS}^3}{\frac{4}{3}\pi R_{core-shell}^3}$$
(3)

212

## 213 3 Results and discussion

#### 214 **3.1 Hygroscopic growth of initially well-mixed aerosol particles**

215 Figure 2 shows the measured hygroscopic growth factors of the AS, PA, and initially well-mixed 216 AS with different mass fractions of PA particles with dry diameter 100nm against RH, respectively. 217 During the hydration mode, there is no change in size until a slow increase occurs at 60 % RH in 218 the case of initially well-mixed AS/PA aerosol particles. This increase may occur because the PA 219 uptakes a small amount of water. However, an abrupt increase in the hygroscopic growth factor is 220 observed at 75 % RH for initially well-mixed particles containing 50 and 75 wt % PA, of which the 221 growth factor is higher than that of pure PA aerosol particles  $(1.09 \pm 0.01 \text{ nm from measurements})$ 222 shown in Fig. 2) at the same RH. An interesting, yet contrasting phenomenon is that water uptake 223 for initially well-mixed particles containing 50 wt % PA components is relatively higher than that 224 of mixtures containing 75 wt % PA at 75% RH. One possible reason is that the full deliquescence 225 of AS in the initially well-mixed particles with 50 wt % PA components is completed at 75 % RH, 226 while AS in the mixtures containing 75 wt % PA components is partially deliquescent. A decrease

227	in the hygroscopic growth factor of initially well-mixed AS/PA particles with increasing mass
228	fractions of PA is observed at RH above 80 %. For example, the measured growth factors for initially
229	well- mixed aerosol particles containing 25, 50, and 75 wt % PA are 1.36, 1.28, and 1.19 at 80 %
230	RH, respectively, lower than the growth factor of 1.45 for pure deliquesced AS particles (value from
231	measurements shown in Fig. 2) at the same RH. Also, the measured hygroscopic growth factors
232	within experimental uncertainty were in good agreement with the results from the initially well-
233	mixed particles performed by Hämeri et al. (2002). In addition, with increasing mass fractions of
234	PA in the initially well-mixed particles, the smoothing of the hygroscopic growth factor curves is
235	obvious, indicating that the PA aerosol particles have a significant effect on the water uptake of
236	initially well-mixed AS/PA particles such as a shift or suppression of DRH of AS in the mixed
237	particles. For example, in the case of 1:3 mixtures of AS:PA (by mass), 75 wt % PA in the initially
238	well-mixed particles suppresses the deliquescence of AS, i.e., AS in the initially well-mixed
239	particles slowly dissolve into the liquid phase due to continuous water uptake of PA prior to the
240	deliquescence relative humidity of AS (80 % RH). This similar phenomenon was observed by
241	previous studies (e.g., Hämeri et al., 2002; Qiu and Zhang, 2013). For example, Qiu and Zhang
242	(2013) observed that mixture particles consisting of dimethylaminium sulfate and AS exhibited a
243	moderate growth by water uptake in the RH range of 40 %-70 % RH. The calculated growth factors
244	from a model based on the ZSR relation agree well with the hygroscopic growth factors of initially
245	well-mixed AS/PA particles when accounting for measurement uncertainty. A possible reason for
246	this good agreement is that the measured growth factors referring to the water uptake contribution
247	by PA in the ZSR relation are obtained from the fitted growth curve of pure PA particles (as shown
248	in Fig. 2. Fitted expression, Eq. (1)). Thus, in the case of initially well-mixed AS/PA, relatively good

249	agreement with the experimental growth factors of mixtures with 25, 50, and 75 wt % PA
250	demonstrates that individual components independently absorb water in proportion to their volume.
251	However, the discrepancy between measured growth factor of initially well-mixed AS/PA particles
252	at 75 % RH and the predicted growth factors by using the ZSR relation may be due to the molecule
253	interaction between organic molecular and completely or partially dissolved AS ions. A similar
254	phenomenon was reported for well-mixed mixtures of AS + levoglucosan in the previous study by
255	Lei et al (2014, 2018). In addition, the use of the AIOMFAC-LLE model to predict LLPS and RH-
256	dependent water content of AS/PA particles containing 46 wt % PA under hydration conditions. As
257	shown in Fig. S3. At low RH, AS forms a crystalline phase prior to its deliquescence. Thus, AS is
258	predominantly partitioned to the solid phase ( $\delta$ ), while PA is found in a separate amorphous/liquid
259	phase that further contains water (in solid-liquid equilibrium) up to the complete AS deliquescence
260	at ~79 % RH. Two liquid phases, one AS-rich ( $\alpha$ ) and one PA-rich ( $\beta$ ), are predicted to coexist
261	between 79 % $<$ RH $<$ 96 % in the hydration case. At RH above 96 %, a single liquid phase is the
262	stable state. Although, we note that AIOMFAC-LLE may have an uncertainty of several % RH in
263	terms of the RH range in which LLPS is predicted (Song et al., 2012).

## 264 **3.2** Hygroscopic growth of core-shell-generated structured aerosol particles

Figure 3 shows the measured hygroscopic growth of core-shell-generated structured particles as a function of RH. Here, we investigated the hygroscopic behavior of samples of various AS core particle sizes (AS particle dry diameter of 100, 150, and 200nm) and coating (PA coating of 10, 20, 30, and 50nm), respectively. The core-shell-generated structured particles start to absorb a small amount of water at RH lower than the DRH of AS due to the organic coating. A similar behavior has been observed for core-shell-generated structured particles containing AS and palmitic acid by

271	Garland et al. (2005), where early water uptake and reduced hygroscopic growth after deliquescence
272	of AS (compared to pure AS aerosols) were reported. A reduction of the hygroscopic growth factors
273	of core-shell-generated particles becomes obvious as the thickness of the PA shell increases after
274	the deliquescence of core-shell-generated particles. For example, the measured growth factor value
275	at 80% RH is 1.45, 1.40, 1.32, and 1.28 for core-shell-generated particles containing 100 nm AS
276	and 10, 20, 30, and 50 nm coating PA shell, respectively. The kinetic limitation on the core-shell-
277	generated particles is expected to increase considerably with increasing the thickness of the coating
278	PA shells, which retards the transport rate of water molecules across core-shell-generated aerosol
279	particles/air interface. In addition, the measured hygroscopic growth factor of core-shell-generated
280	AS/PA mixtures is predicted by the ZSR relation, which is based on the hygroscopic growth factors
281	of AS and PA derived from the E-AIM predictions for AS and the fitted GF curve (Eq. 1). The ZSR-
282	based predictions are lower than that of core-shell-generated aerosol particles at RH in the range of
283	5 % - 90 %. The underprediction of the ZSR relation was also observed in the literature (Chan et al.,
284	2006; Sjogren et al., 2007). Sjogren et al. (2006) observed an enhanced water uptake of mixtures
285	consisting of AS and adipic acid with different mass ratios (1:2, 1:3, and 1:4) at RH above 80 %
286	compared with ZSR relation in the hydration condition. They assumed that adipic acid is more likely
287	to enclose the water-soluble AS in veins and cavities, which results in easy uptake of water and a
288	negative curvature of the solution meniscus at the opening of the vein compared to a flat or convex
289	particle surface. Therefore, in the case of AS/PA core-shell-generated particles, one potential reason
290	for the underestimation of the measured growth factor by ZSR relation is the morphology effect on
291	the core-shell-generated structured AS/PA particle. To be specific, due to the potential presence of
292	polycrystalline AS, containing pores, cracks, and veins (Zelenyuk et al., 2006; Sjogren et al., 2007),

when RH approaches 80 %, these pores or veins may fill with aqueous PA solution. Thus, water molecules may be easier to diffuse to the veins or pores than particle surface. At RH above 80 %, deliquesced AS is more likely to mixed partial aqueous solution PA. The resulting effect of the arrangement and restructuring of core-shell structured particles may change the hygroscopicity, morphology, and mixing state of the core-shell-generated particles (Chan et al., 2006; Sjogren et al., 2007).

299 Figure 4 shows that the experimental water absorption of the varying size of AS core coated with 300 50-nm PA shell in the hydration condition. In the case of 50-nm PA shell coated with a certain size 301 of the AS core (100, 150, and 200 nm) with respect to 68, 55, and 46 wt % PA in the core-shell-302 generated particles, It exhibits an increase in hygroscopic growth factor of core-shell-generated 303 particles at RH below 80 % as the size of AS core decreases. However, a decrease in hygroscopic 304 growth factor of core-shell-generated mixtures is observed at RH above 80 % with decreasing the 305 size of the AS core. This indicates that the 50-nm PA shell in the core-shell-generated particles have 306 predominantly contributed to the hygroscopic growth of core-shell-generated particles at low RH. 307 At high RH (e.g., after AS deliquescence), however, 50-nm PA coating shows a weak kinetic 308 limitations for water uptake by core-shell-generated particles as the size of AS core increases. For 309 example, the measured growth factor values are 1.28, 1.34, and 1.40 at 80 % RH for 100-200 nm 310 AS core in the mixed particles, respectively. The discrepancy between measured hygroscopic 311 growth factors and predicted hygroscopic growth factors of core-shell-generated particles by ZSR 312 relation, as discussed in Sect. 3.2, is due to the morphology effect. For ZSR prediction, it assumes 313 that volume fraction of AS components is constant according to the ratio of the volume of AS core 314 in the sphere to the volume of a core-shell-generated sphere based on Eq. (3). Without considering

315 morphology effect, the ZSR prediction results in an underestimation of hygroscopic growth factors316 of core-shell-generated particles.

# 317 3.3 Comparison of core-shell-generated and initially well-mixed AS/PA aerosol particles Figure 5 shows the hydration curves of different AS cores coated with the different mass fractions 318 319 of PA loading (shown in the Supporting Information Table S2) in comparison with those of the 320 initially well-mixed with the same PA mass fractions particles, pure AS particles, and pure PA particles in the range of 5 % - 90 % RH. The effect of the coating PA on core-shell-generated 321 322 particles becomes more pronounced than that of PA in the initially well-mixed particles at RH below 323 70 % as shown in Fig 5a-b, which leads to higher amounts of water absorption at low RH. However, 324 compared to Fig 5a-b, Fig. 5c shows the hygroscopic growth factors of initially well-mixed AS/PA 325 is slightly higher than that of AS/PA core-shell-generated particles with 46 wt % PA. At 75 % RH, 326 the measured growth factor value of core-shell-generated particles is lower than that of internally 327 well-mixed mixtures in the PA mass fraction range from 68 to 46 wt % due to the mass transfer 328 limitations of water vapor transport to the AS core in the core-shell-generated particles. For the 329 initially well-mixed AS/PA particles, however, partial dissolution of AS into the liquid PA phase 330 may lead to more water uptake by initially well-mixed particles. For example, for the core-shell-331 generated mixtures with 68 wt % PA loading, the experimental growth factor value is 1.09 at 75 % 332 RH, relative to the growth factor of 1.17 of initially well-mixed mixtures AS/PA. After an abrupt 333 increase in particle diameter of mixed particles, the core-shell-generated AS/PA particles uptake 334 slightly more water than initially well-mixed AS/PA with the same mass fractions of PA as RH 335 increases above 80 %. Accordingly, at high RH, the occurrence of microscopical restructuring of 336 core-shell-generated particles may affect their size. A similar hygroscopic behavior was observed in

337 previous papers (Chan et al., 2006; Sjogren et al., 2007; Maskey et al., 2014). Chan et al. (2006) 338 investigated hygroscopicity of 49 wt % glutaric acid coated on AS core during two continuous hydration cycles: the experimental growth factor of the fresh core-shell of AS and glutaric acid in 339 340 the first hydration cycle is slightly higher than those in the second hydration cycle with the same 341 mass fractions of glutaric acid. They suggested that the mixing state has changed from core-shell to 342 well-mixed state during the humidification process. Also, a slightly higher growth factor of core-343 shell particles than that of well-mixed particles was found when comparing the hygroscopic growth 344 factors of 49 wt % glutaric acid coated on AS core with that of well-mixed mixtures of AS with the 345 same mass fractions of glutaric acid from different papers (Choi et al., 2002; Chan et al., 2006). 346 However, a contrasting phenomenon was observed in the previous study (Maskey et al., 2014). 347 Maskey et al. (2014) investigated the hygroscopic behavior of the internal mixtures consisting AS 348 coated with either succinic acid or levoglucosan in the different mixing state with the same volume 349 fractions of organic compounds. The growth factor of core-shell particles consisting of AS and 350 succinic acid is lower than that of the well-mixed particles, while experimental values for core-shell 351 of AS/levoglucosan particles are close to those of the well-mixed mixtures. The possible reasons for 352 the difference between our study and results from Maskey et al. (2014) are physical properties of 353 the organic components, such as hygrosocpicity, viscosity, volatility, and water uptake coefficients. 354 Therefore, different kinds of organic compounds have a different effect in the hygroscopic growth 355 of mixtures, including the core-shell and the well-mixed state. For example, no hygroscopic growth 356 was observed up to 99 % RH for pure succinic acid particles (shown in Fig. S1a). Peng et al. (2001) 357 measured the DRH of succinic acid at 99 % RH using a bulk solution at 24 °C. Also, Henning et al. 358 (2002) observed no hygroscopic growth of soot/succinic acid core-shell particles in the hydration

359	mode using HTDMA. In the case of AS/succinic acid core-shell particles, No water uptake by AS
360	coated succinic acid shell was observed before 80 % RH, while there is a gradual increase in water
361	absorption of core-shell-generated particles prior to the deliquescence of AS with different mass
362	fraction of PA components as shown in Fig 5a-c. This suggested the physical state of shell is solid
363	and liquid for Maskey et al. (2014) and our measurements, respectively. At RH above 80 %, the
364	kinetic limitation on the water vapor uptake through solid shell into the core is more obvious than
365	that through liquid shell into the core (i.e., liquid diffusion coefficient of water vapor is the range of
366	$10^{-10}$ to $10^{-9}$ m <sup>2</sup> s <sup>-1</sup> , solid diffusion coefficient of water vapor is the range of $10^{-13}$ - $10^{-14}$ m <sup>2</sup> s <sup>-1</sup> at 25°C).
367	This can lead to different hygroscopic behavior of core-shell particles. In the case of
368	AS/levoglucosan measured by Maskey et al. (2014), they found that the slightly higher growth
369	factors for well-mixed particles are than core-shell aerosol particles (88-nm AS core coated by 12-
370	nm levoglucosan). The mass fraction of levoglucosan in the core-shell-generated particles is $\sim 29$
371	wt %. In our study, AS coated with PA with mass fraction range is between 46 to 68 wt %. Next,
372	using the low mass fraction of PA (e.g., 29 wt %) in core-shell-generated aerosol particles is to be
373	investigated. In addition, for the AS/PA mixture aerosol particles containing 46-68 wt % PA, the
374	measured growth factors of initially well-mixed AS/PA particles are in good agreement with the
375	ZSR relation prediction comparing with that of core-shell-generated AS/PA particles.

- 376
- 377 **4 Summary and conclusion**

378 In this study, we focused on PA to represent common organic compounds produced by various 379 sources, (e.g., vehicles, biomass burning, photo-oxidation). It is found that PA aerosol particles 380 uptake water continuously as RH increases. We further investigated the effect of PA coating on the

381	hygroscopicity of core-shell-generated aerosol particles. As PA coating thickness increases, the
382	hygroscopic growth factor of AS/PA core-shell-generated particles increases prior to the
383	deliquescence of AS, but the water uptake decreases at RH above 80 %. Furthermore, we compared
384	the hygroscopic behavior of AS/PA core-shell-generated particles with that of AS/PA initially well-
385	mixed particles. Due to mixing state effects, higher hygroscopic growth factors of AS/PA core-shell-
386	generated particles, compared to that of initially well-mixed particles, were observed in this study
387	at RH above 80 %. In addition, the ZSR relation prediction is in good agreement with measured
388	results of AS/PA initially well-mixed particles, but leads to the underestimation of the hygroscopic
389	growth factor of AS/PA core-shell-generated particles at RH above 80 %. We attribute these
390	discrepancies to the morphology effect when AS deliquesces in the core-shell-generated particles.
391	There are a vast number of internally mixed organic-inorganic aerosol particles existing in the
392	atmosphere. The hygroscopicity behavior of mixture particles exhibits variability during RH cycles
393	depending on the chemical composition, size, and mixing state. Humidity cycles may lead to liquid-
394	liquid phase separation, e.g., in the form of core-shell aerosol particles, including at higher RH or
395	in the salt-supersaturated concentration range. Also, due to the different physicochemical properties
396	of organic compounds (e.g., viscosity, solubility, physical state, and morphology), the equilibrium
397	time varies for these organic coated with inorganic aerosol particles. Therefore, potential kinetic
398	limitations in the HTDMA-measured hygroscopicity of core-shell aerosol particles is to be
399	investigated in both humidification and dehumidification conditions.

# 401 Data availability

402 Readers who are interested in the data should contact Ting Lei (<u>ting.lei@mpic.de</u>)

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- 406 Author contributions: W.G.W designed and led the study. W.G.W and T.L assembled the
- 407 coating-HTDMA. T.L performed the experiments and then wrote the manuscript. W.G.W supported

408 the experiments. All co-authors discussed the results and commented on manuscript.

409

### 410 **Reference**

- 411 Abo Riziq, A., A., Trainic, M., Erlick, C., Segre, E., and Rudich, Y.: Extinction efficiencies of coated
- 412 absorbing aerosols measured by cavity ring down aerosol spectrometry, Atmos. Chem. Phys., 8,
  413 1823-1833, 2008.
- 414 Abo Riziq, A., Erlick, C., Dinar, E., and Rudich, Y.: Optical properties of absorbing and non-
- 415 absorbing aerosols retrieved by cavity ring down (CRD) spectroscopy, Atmos. Chem. Phys., 7,
- 416 1523-1536, 2007.
- 417 Ansari, A. S. and Pandis, S. N.: Water Absorption by Secondary Organic Aerosol and Its Effect on
- 418 Inorganic Aerosol Behavior, Environmental Science & Technology, 34, 71-77, 2000.
- 419 Bertram, A. K., Martin, S. T., Hanna, S. J., Smith, M. L., Bodsworth, A., Chen, Q., Kuwata, M., Liu,
- 420 A., You, Y., and Zorn, S. R.: Predicting the relative humidities of liquid-liquid phase separation,
- 421 efflorescence, and deliquescence of mixed particles of ammonium sulfate, organic material, and
- 422 water using the organic-to-sulfate mass ratio of the particle and the oxygen-to-carbon elemental
- ratio of the organic component, Atmos. Chem. Phys., 11, 10995-11006, 2011.
- 424 Brooks, S. D., DeMott, P. J., and Kreidenweis, S. M.: Water uptake by particles containing humic

- 425 materials and mixtures of humic materials with ammonium sulfate, Atmospheric Environment, 38,
- 426 1859-1868, 2004.
- 427 Chan, M. N., Lee, A. K. Y., and Chan, C. K.: Responses of ammonium sulfate particles coated with
- 428 glutaric acid to cyclic changes in relative humidity: Hygroscopicity and Raman characterization,
- 429 Environmental Science & Technology, 40, 6983-6989, 2006.
- 430 Cheng, Y. F., Wiedensohler, A., Eichler, H., Heintzenberg, J., Tesche, M., Ansmann, A., Wendisch,
- 431 M., Su, H., Althausen, D., Herrmann, H., Gnauk, T., Brüggemann, E., Hu, M., and Zhang, Y. H.:
- 432 Relative humidity dependence of aerosol optical properties and direct radiative forcing in the surface
- 433 boundary layer at Xinken in Pearl River Delta of China: An observation based numerical study,
- 434 Atmospheric Environment, 42, 6373-6397, 2008.
- 435 Choi, M. Y. and Chan, C. K.: The Effects of Organic Species on the Hygroscopic Behaviors of
- 436 Inorganic Aerosols, Environmental Science & Technology, 36, 2422-2428, 2002.
- 437 Ciobanu, V. G., Marcolli, C., Krieger, U. K., Weers, U., and Peter, T.: Liquid-Liquid Phase
- 438 Separation in Mixed Organic/Inorganic Aerosol Particles, The Journal of Physical Chemistry A, 113,
- 439 10966-10978, 2009.
- 440 Cruz, C. N. and Pandis, S. N.: Deliquescence and Hygroscopic Growth of Mixed Inorganic–Organic
- 441 Atmospheric Aerosol, Environmental Science & Technology, 34, 4313-4319, 2000.
- 442 Eichler, H., Cheng, Y. F., Birmili, W., Nowak, A., Wiedensohler, A., Brüggemann, E., Gnauk, T.,
- 443 Herrmann, H., Althausen, D., Ansmann, A., Engelmann, R., Tesche, M., Wendisch, M., Zhang, Y.
- 444 H., Hu, M., Liu, S., and Zeng, L. M.: Hygroscopic properties and extinction of aerosol particles at
- ambient relative humidity in South-Eastern China, Atmospheric Environment, 42, 6321-6334, 2008.
- 446 Falkovich, A. H., Schkolnik, G., Ganor, E., and Rudich, Y.: Adsorption of organic compounds

- 447 pertinent to urban environments onto mineral dust particles, Journal of Geophysical Research:
- 448 Atmospheres, 109, n/a-n/a, 2004.
- 449 Ganguly, D., Jayaraman, A., Rajesh, T. A., and Gadhavi, H.: Wintertime aerosol properties during
- 450 foggy and nonfoggy days over urban center Delhi and their implications for shortwave radiative
- 451 forcing, Journal of Geophysical Research: Atmospheres, 111, 2006.
- 452 Garland, R. M., Wise, M. E., Beaver, M. R., DeWitt, H. L., Aiken, A. C., Jimenez, J. L., and Tolbert,
- 453 M. A.: Impact of palmitic acid coating on the water uptake and loss of ammonium sulfate particles,
- 454 Atmos. Chem. Phys., 5, 1951-1961, 2005.
- 455 Gupta, D., Kim, H., Park, G., Li, X., Eom, H. J., and Ro, C. U.: Hygroscopic properties of NaCl and
- 456 NaNO<sub>3</sub> mixture particles as reacted inorganic sea-salt aerosol surrogates, Atmos.
- 457 Chem. Phys., 15, 3379-3393, 2015.
- 458 Hämeri, K., Charlson, R., and Hansson, H.-C.: Hygroscopic properties of mixed ammonium sulfate
- and carboxylic acids particles, AIChE Journal, 48, 1309-1316, 2002.
- 460 Heintzenberg, J., Maßling, A., and Birmili, W.: The connection between hygroscopic and optical
- 461 particle properties in the atmospheric aerosol, Geophysical Research Letters, 28, 3649-3651, 2001.
- 462 Hodas, N., Zuend, A., Mui, W., Flagan, R., and Seinfeld, J.: Influence of particle-phase state on the
- 463 hygroscopic behavior of mixed organic-inorganic aerosols, Atmospheric Chemistry and Physics,
- 464 15, 5027-5045, 2015.
- 465 Hori, M., Ohta, S., Murao, N., and Yamagata, S.: Activation capability of water soluble organic
- 466 substances as CCN, Journal of Aerosol Science, 34, 419-448, 2003.
- 467 Huff Hartz, K. E., Tischuk, J. E., Chan, M. N., Chan, C. K., Donahue, N. M., and Pandis, S. N.:
- 468 Cloud condensation nuclei activation of limited solubility organic aerosol, Atmospheric

- 469 Environment, 40, 605-617, 2006.
- 470 Jing, B., Tong, S., Liu, Q., Li, K., Wang, W., Zhang, Y., and Ge, M.: Hygroscopic behavior of
- 471 multicomponent organic aerosols and their internal mixtures with ammonium sulfate, Atmospheric
- 472 Chemistry and Physics, 16, 4101-4118, 2016.
- 473 K., C. M. N. a. C. C.: Mass transfer effects in hygroscopic measurements of aerosol particles,
- 474 Atmospheric Chemistry and Physics, 2005. 2005.
- 475 Kawamura, K. and Ikushima, K.: Seasonal changes in the distribution of dicarboxylic acids in the
- 476 urban atmosphere, Environmental Science & Technology, 27, 2227-2235, 1993.
- 477 Kleindienst, T. E., Smith, D. F., Li, W., Edney, E. O., Driscoll, D. J., Speer, R. E., and Weathers, W.
- 478 S.: Secondary organic aerosol formation from the oxidation of aromatic hydrocarbons in the
- 479 presence of dry submicron ammonium sulfate aerosol, Atmos Environ, 33, 3669-3681, 1999.
- 480 Kreidenweis, S. M., Koehler, K., DeMott, P. J., Prenni, A. J., Carrico, C., and Ervens, B.: Water
- 481 activity and activation diameters from hygroscopicity data Part I: Theory and application to
- 482 inorganic salts, Atmospheric Chemistry and Physics, 5, 1357-1370, 2005.
- 483 Lang-Yona, N., Abo-Riziq, A., Erlick, C., Segre, E., Trainic, M., and Rudich, Y.: Interaction of
- 484 internally mixed aerosols with light, Physical Chemistry Chemical Physics, 12, 21-31, 2010.
- 485 Lei, T., Zuend, A., Cheng, Y., Su, H., Wang, W., and Ge, M.: Hygroscopicity of organic surrogate
- 486 compounds from biomass burning and their effect on the efflorescence of ammonium
- 487 sulfate in mixed aerosol particles, Atmos. Chem. Phys., 18, 1045-1064, 2018.
- 488 Lei, T., Zuend, A., Wang, W. G., Zhang, Y. H., and Ge, M. F.: Hygroscopicity of organic compounds
- 489 from biomass burning and their influence on the water uptake of mixed organic ammonium sulfate
- 490 aerosols, Atmospheric Chemistry and Physics, 14, 1-20, 2014.

- 491 Lesins, G., Chylek, P., and Lohmann, U.: A study of internal and external mixing scenarios and its
- 492 effect on aerosol optical properties and direct radiative forcing, Journal of Geophysical Research:
- 493 Atmospheres, 107, AAC 5-1-AAC 5-12, 2002.
- Liu, Q., Jing, B., Peng, C., Tong, S., Wang, W., and Ge, M.: Hygroscopicity of internally mixed
- 495 multi-component aerosol particles of atmospheric relevance, Atmospheric Environment, 125, 69496 77, 2016.
- 497 Liu, X., Gu, J., Li, Y., Cheng, Y., Qu, Y., Han, T., Wang, J., Tian, H., Chen, J., and Zhang, Y.: Increase
- 498 of aerosol scattering by hygroscopic growth: Observation, modeling, and implications on visibility,
- 499 Atmospheric Research, 132, 91-101, 2013.
- 500 Liu, X., Zhang, Y., Cheng, Y., Hu, M., and Han, T.: Aerosol hygroscopicity and its impact on
- atmospheric visibility and radiative forcing in Guangzhou during the 2006 PRIDE-PRD campaign,
- 502 Atmospheric Environment, 60, 59-67, 2012.
- 503 Martin, A. C., Cornwell, G. C., Atwood, S. A., Moore, K. A., Rothfuss, N. E., Taylor, H., DeMott,
- 504 P. J., Kreidenweis, S. M., Petters, M. D., and Prather, K. A.: Transport of pollution to a remote
- 505 coastal site during gap flow from California's interior: impacts on aerosol composition, clouds, and
- 506 radiative balance, Atmos. Chem. Phys., 17, 1491-1509, 2017.
- 507 Maskey, S., Chong, K. Y., Kim, G., Kim, J.-S., Ali, A., and Park, K.: Effect of mixing structure on
- 508 the hygroscopic behavior of ultrafine ammonium sulfate particles mixed with succinic acid and
- 509 levoglucosan, Particuology, 13, 27-34, 2014.
- 510 Miñambres, L., Sánchez, M. N., Castaño, F., and Basterretxea, F. J.: Hygroscopic Properties of
- 511 Internally Mixed Particles of Ammonium Sulfate and Succinic Acid Studied by Infrared
- 512 Spectroscopy, The Journal of Physical Chemistry A, 114, 6124-6130, 2010.

- 513 Pagels, J., Khalizov, A. F., McMurry, P. H., and Zhang, R. Y.: Processing of Soot by Controlled
- 514 Sulphuric Acid and Water Condensation—Mass and Mobility Relationship, Aerosol Science and
- 515 Technology, 43, 629-640, 2009.
- 516 Peng, C., Chan, M. N., and Chan, C. K.: The Hygroscopic Properties of Dicarboxylic and
- 517 Multifunctional Acids: Measurements and UNIFAC Predictions, Environmental Science &
- 518 Technology, 35, 4495-4501, 2001.
- 519 Qiu, C. and Zhang, R.: Multiphase chemistry of atmospheric amines, Physical Chemistry Chemical
- 520 Physics, 15, 5738-5752, 2013.
- 521 Rogge, W. F., Mazurek, M. A., Hildemann, L. M., Cass, G. R., and Simoneit, B. R. T.: Quantification
- 522 of urban organic aerosols at a molecular level: Identification, abundance and seasonal variation,
- 523 Atmospheric Environment. Part A. General Topics, 27, 1309-1330, 1993.
- 524 Rose, D., Gunthe, S. S., Su, H., Garland, R. M., Yang, H., Berghof, M., Cheng, Y. F., Wehner, B.,
- 525 Achtert, P., Nowak, A., Wiedensohler, A., Takegawa, N., Kondo, Y., Hu, M., Zhang, Y., Andreae, M.
- 526 O., and Poschl, U.: Cloud condensation nuclei in polluted air and biomass burning smoke near the
- 527 mega-city Guangzhou, China -Part 2: Size-resolved aerosol chemical composition, diurnal cycles,
- and externally mixed weakly CCN-active soot particles, Atmospheric Chemistry and Physics, 11,
- 529 2817-2836, 2011.
- 530 Rudich, Y.: Laboratory Perspectives on the Chemical Transformations of Organic Matter in
- 531 Atmospheric Particles, Chemical Reviews, 103, 5097-5124, 2003.
- 532 Saxena, P., Hildemann, L. M., McMurry, P. H., and Seinfeld, J. H.: Organics alter hygroscopic
- behavior of atmospheric particles, Journal of Geophysical Research: Atmospheres, 100, 18755-
- 534 18770, 1995.

- 535 Schauer, J. J. and Cass, G. R.: Source Apportionment of Wintertime Gas-Phase and Particle-Phase
- 536 Air Pollutants Using Organic Compounds as Tracers, Environmental Science & Technology, 34,

537 1821-1832, 2000.

- 538 Schauer, J. J., Fraser, M. P., Cass, G. R., and Simoneit, B. R. T.: Source Reconciliation of
- 539 Atmospheric Gas-Phase and Particle-Phase Pollutants during a Severe Photochemical Smog
- 540 Episode, Environmental Science & Technology, 36, 3806-3814, 2002.
- 541 Schauer, J. J., Rogge, W. F., Hildemann, L. M., Mazurek, M. A., Cass, G. R., and Simoneit, B. R.
- 542 T.: Source apportionment of airborne particulate matter using organic compounds as tracers,
- 543 Atmospheric Environment, 30, 3837-3855, 1996.
- 544 Schwarz, J. P., Gao, R. S., Fahey, D. W., Thomson, D. S., Watts, L. A., Wilson, J. C., Reeves, J. M.,
- 545 Darbeheshti, M., Baumgardner, D. G., Kok, G. L., Chung, S. H., Schulz, M., Hendricks, J., Lauer,
- 546 A., Kärcher, B., Slowik, J. G., Rosenlof, K. H., Thompson, T. L., Langford, A. O., Loewenstein, M.,
- and Aikin, K. C.: Single-particle measurements of midlatitude black carbon and light-scattering
- 548 aerosols from the boundary layer to the lower stratosphere, Journal of Geophysical Research:
- 549 Atmospheres, 111, n/a-n/a, 2006.
- 550 Shi, Y. J., Ge, M. F., and Wang, W. G.: Hygroscopicity of internally mixed aerosol particles
- 551 containing benzoic acid and inorganic salts, Atmospheric Environment, 60, 9-17, 2012.
- 552 Shiraiwa, M., Zuend, A., Bertram, A. K., and Seinfeld, J. H.: Gas-particle partitioning of
- atmospheric aerosols: interplay of physical state, non-ideal mixing and morphology, Physical
- 554 Chemistry Chemical Physics, 15, 11441-11453, 2013.
- 555 Sjogren, S., Gysel, M., Weingartner, E., Baltensperger, U., Cubison, M. J., Coe, H., Zardini, A. A.,
- 556 Marcolli, C., Krieger, U. K., and Peter, T.: Hygroscopic growth and water uptake kinetics of two-

- 557 phase aerosol particles consisting of ammonium sulfate, adipic and humic acid mixtures, Journal of
- 558 Aerosol Science, 38, 157-171, 2007.
- 559 Song, M., Ham, S., Andrews, R. J., You, Y., and Bertram, A. K.: Liquid–liquid phase separation in
- 560 organic particles containing one and two organic species: importance of the average
- 561 O : C, Atmos. Chem. Phys., 18, 12075-12084, 2018.
- 562 Song, M., Marcolli, C., Krieger, U. K., Zuend, A., and Peter, T.: Liquid-liquid phase separation and
- 563 morphology of internally mixed dicarboxylic acids/ammonium sulfate/water particles, Atmospheric
- 564 Chemistry and Physics, 12, 2691-2712, 2012a.
- 565 Song, M., Marcolli, C., Krieger, U. K., Zuend, A., and Peter, T.: Liquid-liquid phase separation in
- 566 aerosol particles: Dependence on O:C, organic functionalities, and compositional complexity,
- 567 Geophysical Research Letters, 39, 2012b.
- 568 Spindler, C., Riziq, A. A., and Rudich, Y.: Retrieval of Aerosol Complex Refractive Index by
- 569 Combining Cavity Ring Down Aerosol Spectrometer Measurements with Full Size Distribution
- 570 Information, Aerosol Science and Technology, 41, 1011-1017, 2007.
- 571 Stock, M., Cheng, Y. F., Birmili, W., Massling, A., Wehner, B., Müller, T., Leinert, S., Kalivitis, N.,
- 572 Mihalopoulos, N., and Wiedensohler, A.: Hygroscopic properties of atmospheric aerosol particles
- 573 over the Eastern Mediterranean: implications for regional direct radiative forcing under clean and
- polluted conditions, Atmos. Chem. Phys., 11, 4251-4271, 2011.
- 575 Su, H., Rose, D., Cheng, Y. F., Gunthe, S. S., Massling, A., Stock, M., Wiedensohler, A., Andreae,
- 576 M. O., and Pöschl, U.: Hygroscopicity distribution concept for measurement data analysis and
- 577 modeling of aerosol particle mixing state with regard to hygroscopic growth and CCN activation,
- 578 Atmos. Chem. Phys., 10, 7489-7503, 2010.

- 579 Tang, I. N. and Munkelwitz, H. R.: Water activities, densities, and refractive indices of aqueous
- 580 sulfates and sodium nitrate droplets of atmospheric importance, Journal of Geophysical Research:
- 581 Atmospheres, 99, 18801-18808, 1994.
- 582 Tie, X., Huang, R.-J., Cao, J., Zhang, Q., Cheng, Y., Su, H., Chang, D., Pöschl, U., Hoffmann, T.,
- 583 Dusek, U., Li, G., Worsnop, D. R., and O'Dowd, C. D.: Severe Pollution in China Amplified by
- 584 Atmospheric Moisture, Scientific Reports, 7, 15760, 2017.
- 585 Wang, G., Kawamura, K., Xie, M., Hu, S., Li, J., Zhou, B., Cao, J., and An, Z.: Selected water-
- soluble organic compounds found in size-resolved aerosols collected from urban, mountain and
- 587 marine atmospheres over East Asia, Tellus B, 63, 371-381, 2011.
- 588 Wex, H., Hennig, T., Salma, I., Ocskay, R., Kiselev, A., Henning, S., Massling, A., Wiedensohler,
- 589 A., and Stratmann, F.: Hygroscopic growth and measured and modeled critical super-saturations of
- an atmospheric HULIS sample, Geophysical Research Letters, 34, n/a-n/a, 2007.
- 591 Xue, H., Khalizov, A. F., Wang, L., Zheng, J., and Zhang, R.: Effects of dicarboxylic acid coating
- on the optical properties of soot, Physical Chemistry Chemical Physics, 11, 7869-7875, 2009.
- 593 You, Y., Smith, M. L., Song, M., Martin, S. T., and Bertram, A. K.: Liquid–liquid phase separation
- 594 in atmospherically relevant particles consisting of organic species and inorganic salts, International
- 595 Reviews in Physical Chemistry, 33, 43-77, 2014.
- 596 Zamora, I. R., Tabazadeh, A., Golden, D. M., and Jacobson, M. Z.: Hygroscopic growth of common
- 597 organic aerosol solutes, including humic substances, as derived from water activity measurements,
- 598 Journal of Geophysical Research: Atmospheres, 116, n/a-n/a, 2011.
- 599 Zawadowicz, M. A., Proud, S. R., Seppalainen, S. S., and Cziczo, D. J.: Hygroscopic and phase
- 600 separation properties of ammonium sulfate/organics/water ternary solutions, Atmos. Chem. Phys.,

601 15, 8975-8986, 2015.

- 602 Zelenyuk, A., Cai, Y., and Imre, D.: From Agglomerates of Spheres to Irregularly Shaped Particles:
- 603 Determination of Dynamic Shape Factors from Measurements of Mobility and Vacuum
- Aerodynamic Diameters, Aerosol Science and Technology, 40, 197-217, 2006.
- 605 Zhang, Q., Canagaratna, M. R., Jayne, J. T., Worsnop, D. R., and Jimenez, J.-L.: Time- and size-
- resolved chemical composition of submicron particles in Pittsburgh: Implications for aerosol
- sources and processes, Journal of Geophysical Research: Atmospheres, 110, n/a-n/a, 2005.
- 608 Zhang, R., Khalizov, A. F., Pagels, J., Zhang, D., Xue, H., and McMurry, P. H.: Variability in
- morphology, hygroscopicity, and optical properties of soot aerosols during atmospheric processing,
- 610 Proceedings of the National Academy of Sciences, 105, 10291-10296, 2008.
- 611 Zhang, Y.-L., Kawamura, K., Watanabe, T., Hatakeyama, S., Takami, A., and Wang, W.: New
- 612 directions: Need for better understanding of source and formation process of phthalic acid in
- aerosols as inferred from aircraft observations over China, Atmospheric Environment, 140, 147-149,
- 614 2016.
- 615 Zhou, Q., Pang, S.-F., Wang, Y., Ma, J.-B., and Zhang, Y.-H.: Confocal Raman Studies of the
- 616 Evolution of the Physical State of Mixed Phthalic Acid/Ammonium Sulfate Aerosol Droplets and
- 617 the Effect of Substrates, The Journal of Physical Chemistry B, 118, 6198-6205, 2014.
- 618 Zuend, A., Marcolli, C., Luo, B. P., and Peter, T.: A thermodynamic model of mixed organic-
- 619 inorganic aerosols to predict activity coefficients, Atmospheric Chemistry and Physics, 8, 4559620 4593, 2008.
- 621 Zuend, A., Marcolli, C., Peter, T., and Seinfeld, J. H.: Computation of liquid-liquid equilibria and
- 622 phase stabilities: implications for RH-dependent gas/particle partitioning of organic-inorganic

623	aerosols, Atmospheric Chem	istry and Physics,	10, 7795-7820, 2010.
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- 624 Zuend, A. and Seinfeld, J. H.: A practical method for the calculation of liquid–liquid equilibria in
- 625 multicomponent organic-water-electrolyte systems using physicochemical constraints, Fluid Phase
- 626 Equilibria, 337, 201-213, 2013.

# 645 Tables

## 646 **Table 1.** Coefficients of the fitted growth curve parameterization to measured growth factor data

# 647 using Eq. (1)

	Chemical Compound	a	b	c	
	Phthalic Acid	0.083116	0.291473	-0.353544	
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## 665 Figures



Figure 1. Schematic diagram of the coating-hygroscopicity tandem differential mobility analyzer. Here, CO: critical orifice; DMA: differential mobility analyzer; RH1 and RH2 (measured RH sensor) represent the RH of aerosol and humidified flow in the inlet of DMA1 and humidifier, respectively. RH3 (measured by dew point mirror) represent the RH of excess air. Valve B is open and valve A is closed to the homogeneous internally mixed-mode experiment. Valve A is open and Valve B is closed to the coating-mode experiment. Black line: aerosol line; Blue line: sheath line; Green line: Milli-Q water. 



Figure 2. Hygroscopic growth factor for 100 nm (dry diameter, RH < 5 %) aerosol particles containing: ammonium sulfate (AS), phthalic acid (PA), and initially well-mixed mixtures of PA and AS with different mass ratio of AS to PA. In comparison, the E-AIM model, the fitted expression Eq. (1), and the ZSR relation predicted growth factors of AS, PA, and initially well-mixed particles with different mass fractions of PA, respectively. 





Figure 3. Hygroscopic growth factor for core-shell-generated of ammonium sulfate (AS) and phthalic

acid (PA) aerosol particles. In comparison, the ZSR relation predicted growth factor of core-shell-

generated aerosol particles (a) 100-nm AS core (b) 150-nm AS core (c) 200-nm AS core.



Figure 4. Hygroscopic growth factor for 100-200 nm ammonium sulfate (AS) core with coating 50 nm

## 707 phthalic acid (PA). In comparison, the ZSR relation predicted growth factor of core-shell-generated

708 aerosol particles with different AS cores.





Figure 5. Hygroscopic growth factor for core-shell-generated and initially well-mixed aerosol particles

with the same dry mass fractions of phthalic acid (PA) containing: (a): 68 wt % PA, (b): 55 wt % PA, (c):

46 wt % PA. In comparison, the E-AIM model, the fitted expression Eq. (1), and the ZSR relation

- 727 predicted growth factors of ammonium sulfate (AS), PA, and internally well-mixed particles with
- 728 different mass fractions of PA, respectively.

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