



### 1 Atmospheric mercury in the Southern Hemisphere – Part 2: Source 2 apportionment analysis at Cape Point station, South Africa.

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## 13 Abstract

14 Mercury (Hg) contamination is ubiquitous. In order to assess its emissions, transport, 15 atmospheric reactivity, and deposition pathways, Hg monitoring stations have been 16 implemented on a global scale over the past 10-20 years. Despite this significant step forward, 17 the monitoring efforts have been insufficient to fulfill our understanding of Hg cycling in the 18 Southern Hemisphere. While oceans make up 80% of the Southern Hemisphere's surface 19 area, little is known about the effects of oceans on Hg cycling in this region. For instance, in 20 the context of growing interest in effectiveness evaluation of Hg mitigation policies, the relative 21 contribution of anthropogenic and legacy emissions to present-day atmospheric Hg levels is 22 unclear. This paper constitutes Part 2 of the study describing a decade of atmospheric Hg 23 concentrations at Cape Point, South Africa, i.e. the first long-term (> 10 years) observations in 24 the Southern Hemisphere. Building on the trend analysis reported in Part 1, here we combine 25 atmospheric Hg data with a trajectory model to investigate sources and sinks of Hg at Cape 26 Point. We find that the continent is the major sink and the Ocean, especially warm regions, is 27 the major source for Hg.

28 Further, we find that mercury concentrations and trends from long range transport are 29 independent of the source region (e.g. South America, Antarctica) and thus indistinguishable. 30 Therefor, by filtering out air masses from source and sink regions we are able to create a 31 dataset representing a southern hemispheric background Hg concentrations. Based on this 32 dataset we were able to show that the inter-annual variability of Hg concentrations is not driven 33 by changes in atmospheric circulation but rather due to changes in global emissions (gold 34 mining and biomass burning).





## 35 1. Introduction

36 Mercury (Hg) is a toxic pollutant that is ubiquitous in the environment. Due to anthropogenic 37 emissions, the amount of mercury in the atmosphere has increased sevenfold since pre-38 industrial times (Amos et al., 2013). Mercury occurs in the atmosphere as gaseous oxidized 39 mercury (GOM), particle bond mercury (PBM) and predominantly as gaseous elemental 40 mercury (GEM). Because of its atmospheric lifetime of about 1 year, once emitted into the 41 atmosphere, GEM is transported on hemispheric and global scales (Slemr et al., 2018). Since 42 2017 usage and emissions of Hg are regulated under the UN Minamata Convention on 43 Mercury (UNEP, 2013). This UN convention forces its member states to assess the current 44 state of mercury pollution, take actions to reduce mercury emissions, and to evaluate the 45 success of these measures on a regular basis.

46 In order to assess the impact of emission reductions on the system it is necessary to better 47 understand the sources and sinks driving atmospheric mercury cycling. Especially in the 48 southern hemisphere there has been a lack of long-term atmospheric observations that allow 49 to investigate and distinguish long-term trends from the natural variability of atmospheric Hg 50 concentrations. So far, the only long-term observations in the southern hemisphere with 51 measurements over more than 10 years have been and are performed at Cape Point (CPT), 52 South Africa, where Hg has been measured since 1995 (Baker et al., 2002, Slemr et al., 2008). 53 At CPT, for the first ten years (September 1995 to December 2004) Hg concentrations showed 54 a decreasing trend (Slemr et al., 2008, Martin et al., 2017) while Martin et al. (2017) identified 55 an increasing trend for the last ten years (March 2007 to June 2015). Yet the reason for the 56 observed trends is unclear and there was no explanation for the change in sign from a 57 decreasing to an increasing trend.

58 This work is presented in two accompanying papers where the first one (Slemr et al., 59 submitted) focuses on long-term trends in the southern hemisphere over the last ten years 60 based on measurements at CPT and Amsterdam Island (AMS) which is operational since 61 2012. The key finding of that paper is that since 2007 mercury concentrations at CPT seem to 62 have been increasing while no significant trend was found in the 2012 – 2017 period both at 63 CPT and AMS. The upward CPT trend in 2007 – 2017 period seems to be driven by 64 exceptionally low Hg concentrations in 2009 and above average concentrations in 2014.

65 Here, we combine ten years of Hg (2007 – 2016) observations at CPT with calculated hourly 66 backward trajectories in order to investigate sources and sinks for mercury and to quantify the 67 impact of long-term changes in atmospheric circulation patterns on observed Hg 68 concentrations at CPT. The aim of this study is to:

- 69 distinguish between local changes at CPT and hemispheric Hg trends;
- 70 identify source and sink regions for Hg at CPT;
- 71 estimate the natural variability of Hg concentrations at CPT in order to distinguish them
- 72 from other effects such as changing emissions.





73 This paper aims to improve our understanding of mercury cycling in the southern hemisphere.
74 For this, we elaborate on the research question whether concentrations and trends observed
75 at CPT are dominated by local signals or representative for mercury cycling across large parts
76 of the southern hemisphere. Based on backward trajectories and statistical modeling we
77 investigate source and sink regions for mercury observed at Cape Point and the impact of
78 inter-annual variability on atmospheric transport patterns and emissions processes.

# 79 2. Methodology

#### 80 2.1 Observations

81 This study is based on ten years (2007-2016) of continuous gaseous elemental Hg 82 measurements at Cape Point (CPT,  $34^{\circ}21$ 'S,  $18^{\circ}29$ 'E), South Africa. The CPT measurement 83 site is part of the GAW (Global Atmospheric Watch) baseline monitoring observatories of the 84 World Meteorological Organization (WMO). It is located at the southernmost tip of the Cape 85 Peninsula on top of the cliff at an altitude of 230 m (a.s.l.). There are no major local Hg sources 86 and the nearest city, Cape Town, is located 60 km to the north (see Fig. 1). The station is in 87 operation since the 1970ties and, besides Hg, several other pollutants are measured on a 88 regular basis. These include CO<sub>2</sub>, CO, ozone, methane, and radon (<sup>222</sup>Rn) which we use to 99 substantiate the findings on mercury. A detailed description of the CPT station can be found in 90 the accompanying paper (Slemr et al., submitted).



91

- 92 Figure 1: Location of the Cape Point site (CPT, red star), at the southernmost tip of the Cape Peninsula,
- 93 and of known anthropogenic mercury emission sources (in g/km<sup>2</sup>, Global Mercury Assessment 2018 94 emission inventory) in South Africa. This map was made with QGIS.





#### 95 2.2 Modeling

96 GEM measurements (gaseous elemental Hg, the dominant form of Hg in the atmosphere 97 (~95%)) at CPT are performed continuously with a 15 min sampling interval. The GEM 98 measurements were aggregated to hourly averages and for each hourly measurement an 99 ensemble of 5-day backward trajectories was calculated using the HYSPLIT model (Stein et 100 al., 2015) (Fig. 2). For the hourly trajectory ensembles we used different starting altitudes in 101 order to capture the model uncertainty due to the model's initial conditions. The HYSPLIT 102 model was run for ten years (2007 to 2016) using GDAS (Global Data Assimilation System) 103 0.5°x0.5° degree meteorological inputs based on the NCEP/NCAR reanalysis dataset (Kalnay 104 et al., 1996, NOAA, 2004).

105



106 Fig. 2: Origin of air masses influencing the Cape Point site (black dot). Gridded back trajectory 107 frequencies using an orthogonal map projection with hexagonal binning. The tiles represent the number 108 of incidences. 2007-2016 hourly back trajectories were computed using the HYSPLIT model (Stein et al. 109 2015) and the figure was made using the R package openair (Carslaw and Ropkins, 2012).

### 110 2.3 Regionalization

111 The trajectories were categorized into six source regions depending on their travel path (Table 112 1). These categories are:

- 113 Local
- 114Air parcels which traveled less than 100 km absolute distance to CPT over the115last three days are considered to be local air masses.
- 116 Continental
- 117 Air parcels that spend more than 80% of travel time over the African continent.
- 118 Eastern Ocean





- 119Air parcels which did not travel over land and did not go west of 30° E within the120last 4 days.
- 121 South American
- 122 Air parcels which were west of 30°W within the last 4 days.
- 123 Antarctic
- 124 Air parcels which were south of 55° S within the last 4 days.
- 125 Atlantic

126	Air masses which do not fall within the other categories and spend more than
127	80% of the time over the Atlantic Ocean. This category makes up the majority of
128	all trajectories.

129 Naturally, the categorization of air parcels depends on the definition of regions of origin and the 130 travel time chosen for the algorithm. We calculated 5-day backward trajectories and 131 experimented with different cutoff values to determine the source regions of the air masses 132 (Table 1). For this study we chose a cutoff time of 4 days to determine long range transport 133 from Antarctica and South America. However, he choise of cuttoff times of 3 or 5 days did not 134 change the conclusions of our study. This decision is based on tests with different cutoff times 135 and on the fact that the uncertainty of the trajectories grows with travel time. Moreover, air 136 parcels are often a mixture of different source regions (e.g. Atlantic/continental). As an 137 additional test for the calculated categorization we used secondary parameters such as <sup>222</sup>Rn, 138 CO, CO<sub>2</sub>, CH<sub>4</sub>, and O<sub>3</sub>. <sup>222</sup>Rn is a radioactive gas of predominantly terrestrial origin with a half-139 life of 3.8 days. Thus, high <sup>222</sup>Rn concentrations mark air masses which recently passed over 140 the continent such as "Continental" and "Local". Other examples are the distinction of long-141 range transport from South America from Atlantic air masses. Here, we would also expect 142 higher concentrations of other anthropogenic pollutants (e.g. <sup>222</sup>Rn, CH<sub>4</sub>).

Days	Antarctic	S. America	Continent al	Eastern O.	Atlantic	Local
2	1778	150	14580	5930	143478	1760
	(1%)	(< 1%)	(9%)	(4%)	(86%)	(1%)
3	11800	3614	10596	7842	132876	926
	(7%)	(2%)	(6%)	(5%)	(79%)	(1%)
4	26696	12756	7928	7882	111770	550
	(16%)	(8%)	(5%)	(5%)	(67%)	(<1%)
5	39710	22960	5906	6876	91666	370
	(24%)	(14%)	(4%)	(4%)	(55%)	(<1%)

143 Table 1: Impact of travel time cutoff on air parcels source region categorization.





### 144 2.4 Identification of source/sink regions

145 In order to evaluate source and sink regions we calculated the 10<sup>th</sup> and 90<sup>th</sup> percentile of GEM 146 measurements for each season (seasons being defined as three-month intervals: DJF 147 (summer), MAM (autumn), JJA (winter), SON (spring). This seasonal filter proved to be 148 necessary to remove the annual cycle in GEM concentrations driven (among others) by the 149 seasonality of emissions, planetary boundary layer height, and transport patterns. 150 Furthermore, we filtered out mercury depletion events, of unknown origin (Brunke et al. 2010), 151 which were defined as hourly average GEM concentrations of less 0.25 ng/m<sup>3</sup>.

For the source/sink region analysis we interpolated hourly trajectory locations onto a polar stereographic grid centered over the South Pole and calculated the total amount of trajectories traveling through each grid cell over the ten year (2007-2016) time span. We then performed the same procedure for the trajectories of the 10<sup>th</sup> and 90<sup>th</sup> percentile GEM concentrations. By dividing these percentile maps by the total amount of trajectories traveling through each grid rell, we created maps indicating the regional prevalence of high and low GEM concentrations. In the theoretical case of perfectly homogeneous, evenly distributed sources and sinks each grid cell would have a value of 0.1 indicating that 10% of all air parcels in each grid cell belong to the 10% highest/lowest GEM observations. Deviations from this uniform distribution are then interpreted as source regions for high/low GEM concentrations. E.g. a value of 0.2 indicates that twice as many high/low GEM concentrations originate from a given grid cell compared to a uniform distribution.

164 To better distinguish the 10th and 90th percentile plots we chose opposite color schemes for 165 the 10th and 90th percentile plots. In the case of the 90th percentile plots red color indicates 166 source regions for high GEM concentrations (i.e. > 0.1) while blue color indicates the absence 167 of sources in this region (Fig 3a). For the 10th percentile plots blue color indicates sink regions 168 for GEM concentrations (Fig 3b) while red color indicates the absence of sinks. It is important 169 to note that an absence of sources is not equal to the presence of sinks and vice versa. Figure 170 3 gives an example of these plots for air masses attributed to the 'Atlantic' category for <sup>222</sup>Rn 171 measurements. This plot serves as an evaluation of the regionalization algorithm. It can be 172 seen that high <sup>222</sup>Rn concentrations are found only in air masses that traveled over the 173 continent (Fig 3a). Similarly, Figure 3b depicts the fact that no measurements with low <sup>222</sup>Rn 174 concentrations were found in air masses that traveled along the coast line, indicating an impact 175 of anthropogenic sources. Finally, this procedure is sensitive to the total amount of trajectories 176 traveling through a grid cell which leads to low signal to noise ratios in the outskirts of the plot 177 where only few trajectories originate at all. We used a cutoff value of 10 hits and discarded all 178 grid cells with fewer hits but this still leads to a few non-significant hot spots at the outskirts of 179 the domain (e.g. Fig 3a in Antarctica).







181 Figure 3: Distribution map for the 90th percentile highest <sup>222</sup>Rn concentrations (left) and the 10th 182 percentile lowest <sup>222</sup>Rn concentrations (right) measured at Cape Point. Values are the dimensionless 183 prevalence of air parcels of a given concentration percentile ranging from 0 to 1.

### 184 **3. Results**

#### 185 3.1 Source and sink regions

186 To gain an initial overview we investigate 10th and 90th percentile maps for all GEM 187 measurements over the whole period 2007-2016 (Fig. 4). It can be seen that low GEM 188 concentrations originate almost exclusively from air masses which traveled over the continent 189 (Fig 4b). This result is in line with a cluster analysis performed by Venter et al. (2015) *"Air* 190 masses that had passed over the very sparsely populated semi-arid Karoo region, almost 191 directly to the north of CPT GAW, were mostly associated with [...] lower GEM values". It is 192 also consistent with the finding of Slemr et al. (2013) that southern Africa, based on Hg vs 193 <sup>222</sup>Rn correlations, is a net sink region. The reason for this is probably a mixture of near zero 194 emissions in the region and increased dry deposition on the surface.

195 Over the Atlantic Ocean, low GEM concentrations are in line with a uniform distribution with 196 values mostly only slightly below the equilibrium value of 0.1. The exception are air masses 197 that travelled over the ocean east of Cape Point where almost no low concentration GEM 198 measurements originated. Looking at the highest 90th percentile of GEM concentrations, air 199 masses travelling over the ocean show a lower abundance with the exception of a patch east 200 of Cape Point (Fig 4a).

201 The picture becomes clearer when plotting trajectories independently for each of the previously 202 defined regions (Fig. 5). It can be seen that air masses from the eastern ocean sector are the 203 predominant source region of air masses with elevated GEM concentrations (Fig. 5e). In this 204 region the Agulhas Current transports warm water from the Indian Ocean to the Atlantic 205 Ocean. For continental air masses (Fig. 5b) certain source regions can be identified. These 206 coincide with known major Hg emitters (Fig. 1). For air masses representing long range 207 transport (Atlantic, South American, Antarctic) frequency values of the 90th percentile highest 208 GEM concentrations are mostly around 10% indicating no specific sources or sinks in these 209 regions.







210 *F*igure 4: Prevalence of highest (90th percentile) (left) and lowest (10th percentile) (right) GEM 211 concentrations using all hourly trajectories over ten years.



Figure 5: 90th percentile highest GEM concentrations for air masses from six source regions. Red colormeans source regions, blue one absence of emissions in that region.





Looking at the 10th percentile of lowest GEM concentrations, regional and continental air masses can be identified as the single most important sink region (Fig. 6a,b). There are also some continental areas with a high prevalence of low Hg concentrations attributed to the Atlantic sector. These can be interpreted as air parcels with a mixed continental/Atlantic travel path that have been attributed as Atlantic air masses by the algorithm as they did not spend enough time over the continent to be attributed to this sector. Finally, there are no air masses with low GEM concentrations originating from the eastern ocean sector (Fig. 6e). Yet again, there is no clear picture concerning air masses from long range transport.



222

223 Fig. 6: 10th percentile highest Hg concentrations for six source regions. Blue colors indicate sink 224 regions, red color absence of sinks. Low values are almost exclusively linked to continental and local air 225 masses.





 $226\,$  To further investigate the processes behind the observed source and sink regions, we regard

- 227 seasonal trajectory maps (Fig. 7). This analysis reveals, that the high GEM concentrations 228 associated to air masses from the eastern ocean occur mainly during austral spring and
- 229 summer. This indicates that temperature or primary production and a related increase in
- 230 evasion of GEM (i.e., reemission of legacy Hg) from the ocean might explain these
- 231 observations.
- 232



Figure 7: Seasonal breakdown of 90th percentile highest Hg concentrations from the eastern ocean Highest Hg concentrations are mostly associated to Austral spring and summer which primary production as a potential source for oceanic Hg releases.

### 237 3.2 Comparison of regionalized data

238 Annual and monthly averages and medians for each source region of 4-days backward 239 trajectory were calculated for GEM, CO<sub>2</sub>, <sup>222</sup>Rn, CO, CH<sub>4</sub>, and O<sub>3</sub>. Here we compare the annual 240 averages and medians for all species and discuss the implications this comparison provides on 241 regionalization.

242 The highest annual average and median GEM concentrations were found for "Eastern Ocean" 243 in all years except for 2015. The lowest annual GEM concentrations were almost always either 244 "Local" (averages in 2007, 2010, 2013; medians in 2007, 2010, 2013, 2014) or "Continental" 245 (averages in 2008, 2009, 2012, 2016; medians in 2008, 2010, 2013, 2015). Annual average 246 and median GEM concentrations for "South American", "Antarctic" and "Atlantic" lie in the 247 middle in varying order.

248 The highest and second highest <sup>222</sup>Rn annual average and median concentrations were found 249 for "Continental" and "Local", respectively, in all years except in 2007, in which the "Local" 250 concentration is larger than the "Continental" one. This is consistent with the terrestrial origin of 251 <sup>222</sup>Rn, a radioactive gas with a half-life of 3.8 days (Zahorowski et al., 2004). The lowest and 252 second lowest <sup>222</sup>Rn concentrations were found in air masses attributed to "South American" 253 and "Antarctic" (averages in all years except 2009, 2010, and 2011; medians in all years 254 except 2010 and 2011), respectively. In the exceptional years the lowest <sup>222</sup>Rn concentrations 255 were "Antarctic" and the second lowest "South American". The concentrations attributed to 256 "Eastern Ocean" and "Atlantic" lie in the middle, with "Eastern Ocean" concentrations being 257 mostly higher than the "Atlantic" ones (always in medians, in averages all years with exception 258 of 2009 and 2010).

259 The highest and second highest annual CO mixing ratios were found in air masses attributed 260 to "Continental" and "Local", respectively, with a few years with reverse order (averages: 2007-261 2009; medians: 2007 and 2016). The lowest annual CO mixing ratios were found in air masses





262 attributed to "Antarctic" (averages in 2007, 2009, 2011, 2013, 2015; medians in 2007, 2009, 263 2011, and 2014) or "South American" (averages in 2008, 2012, 2014, 2016; medians in 2012 264 and 2016). CO mixing ratios for "Atlantic" and "Eastern Ocean" lie in the middle with varying 265 order.

266 The highest annual average and median  $CH_4$  mixing ratios were observed in air masses 267 attributed almost always to "Local" or "Continental" (averages "Local" in 2007, 2008, 2011, 268 2014, 2016 and "Continental" in 2009, 2010, 2012, 2013, 2016; medians "Local" in 2007, 2010, 269 and 2014 – 2016 and "Continental" in 2008, 2009, and 2011 - 2013). Mixing ratios "Eastern 270 Ocean" were mostly the lowest (averages in 2007, 2009-2012, 2015; medians in 2007, 2009 – 271 2011, 2015) or second lowest (averages in 2008 and 2013; medians in 2008, 2012, 2016). 272 Annual mixing ratios for "South American", "Antarctic" and "Atlantic" lie mostly in the middle 273 with varying order.

274 The highest average and median annual CO2 mixing ratios are mostly attributed to air masses 275 of either "Local" or "Continental" origin with only a few exceptions (average in 2007; median in 276 2012). The lowest annual mixing ratios are mostly either "Antarctic" or "South American" with 277 only a few exceptions (average and median in 2012). Annual averages and medians for 278 "Atlantic" and "Eastern Ocean" are mostly in the middle.

279 O<sub>3</sub> is being photochemically produced in NOx rich air masses, i.e. mostly in air masses of 280 continental origin, and destroyed in NOx poor air masses, i.e. mostly over the remote ocean 281 (Monks et al., 1998; Fischer et al., 2015). As expected, annual "Continental" average and 282 median O<sub>3</sub> mixing ratios are highest or second highest (averages highest in 2007, 2009 – 283 2011, 2013 – 2015, second highest in 2008, 2012, and 2016; medians highest in 2007, 2009, 284 2010, and 2014, second highest in 2008, 2011, 2012, 2015 and 2016). The lowest annual 285 average and median O<sub>3</sub> mixing ratios are attributed to "Eastern Ocean" in all years except 286 2008 and 2014, in which the mixing ratios are second lowest. The second lowest annual 287 average and median O<sub>3</sub> mixing ratios are found in air masses attributed to "Atlantic" (averages 288 and medians in 2007, 2009, 2011, 2013, 2015). An exception are the years 2008 and 2014 in 289 which "Atlantic" provides the lowest mixing ratios.

290 The results reported above apply for regionalization using 4-days backward trajectories. 291 Regionalization with 3 or 5-days backward trajectories provides similar results.

In summary, regionalized average and median annual mixing ratios of CO, CO<sub>2</sub>, CH<sub>4</sub>, and especially concentrations of <sup>222</sup>Rn behave as expected for species of terrestrial origin, i.e. with highest values almost always in air masses attributed to "Continental" or Local" and the lowest ones attributed mostly to "Antarctic" or "South American". "Atlantic" and "Eastern Ocean" tend to lie in the middle. Ozone, although not of terrestrial but of photochemical origin in NOx rich environments, also fits this pattern because its mixing ratios are highest in "Local" or "Continental" air masses where the highest NOx mixing ratios are expected. The lowest O<sub>3</sub> mixing ratios in "Eastern Ocean" and "Atlantic" can be perhaps explained by NOx poor environment as in "Antarctic" and "South American" but larger photochemical destruction due to lower latitude and thus higher solar radiation. GEM, with highest concentrations in air masses attributed to "Eastern Ocean" and the lowest one found either in "Local" or "Continental", behaves just in contrary fashion to all the species mentioned above. Its contrary behaviour pattern clearly shows that its sources are predominantly oceanic and the sinks 305 terrestrial.





#### 306 3.3 Regional trends

307 Regionalised trends for GEM, CO<sub>2</sub>, <sup>222</sup>Rn, CO, CH<sub>4</sub>, and O<sub>3</sub> were calculated from regional 308 monthly averages and medians using least square fit. Months with less than 10 measurements 309 were not considered. The trends of <sup>222</sup>Rn and O<sub>3</sub> are insignificant for all regions. For the 310 remaining species we present only trends significant both in monthly averages and medians. 311 The trend differences are tested for significance by comparison of averages (Kaiser and 312 Gottschalk, 1972) using the slope and its uncertainty as an average and its standard deviation, 313 respectively.

314 CO<sub>2</sub> and CH<sub>4</sub> trends are significant for all regions. The regional CO<sub>2</sub> trends are all significantly 315 different (>99.9%) from each other with the exception of the trends of monthly medians of 316 "Eastern Ocean" and "Atlantic". The highest trend was observed in "Continental" air masses 317 (averages 2.24 ± 0.04 ppm yr<sup>-1</sup>, medians 2.23 ± 0.04 ppm yr<sup>-1</sup>) and the lowest in "Local" 318 (averages 2.07 ± 0.12 ppm yr<sup>-1</sup>, medians 2.10 ± 0.11 ppm yr<sup>-1</sup>), the latter with exceptionally 319 high uncertainty. The sequence of the remaining regions is "Eastern Ocean" > "Antarctic" > 320 "Atlantic" > "South American" for trends from monthly averages and "Antarctic" > "Atlantic" > 321 "Eastern Ocean" > "South American" for trends from monthly medians.

322 The highest CH<sub>4</sub> trends were observed in "Continental" in monthly averages (7.52 ± 0.76 ppb 323 yr<sup>-1</sup>) with "Eastern Ocean" being the second highest, and "Eastern Ocean" in monthly medians 324 (7.34 ± 0.55 ppb yr<sup>-1</sup>) with "Continental" being the second highest. The lowest trends were 325 found to be in "South American" both in monthly averages and medians (averages 6.21 ± 0.54 326 ppb yr<sup>-1</sup>, medians 6.37 ± 0.47 ppb yr<sup>-1</sup>). The trends between these extremes were mostly not 327 significantly different. The only significant CO trend both in monthly averages and medians was 328 found for "South American" air masses (averages -1.30 ±0.58 ppb yr<sup>-1</sup>, medians -1.02 ± 0.42 329 ppb yr<sup>-1</sup>) and, therefore, cannot be compared with trends for other regions.

330 Three source regions provide significant trends for GEM, both in averages and medians. The 331 trends for "Antarctic" and "South American" air masses are comparable (averages "Antarctic" 332 with 14.25  $\pm$  2.97 pg m<sup>-3</sup> yr<sup>-1</sup>, "South American" with 14.88  $\pm$  3.95 pg m<sup>-3</sup> yr<sup>-1</sup>; medians 333 "Antarctic" with 12.72  $\pm$  2.86 pg m<sup>-3</sup> yr<sup>-1</sup>, "South American" with 13.46  $\pm$  3.96 pg m<sup>-3</sup> yr<sup>-1</sup>). They 334 are both almost twice as large as the trend for "Atlantic" air masses (averages 8.69  $\pm$  2.54 pg 335 m<sup>-3</sup> yr<sup>-1</sup>, medians 8.49  $\pm$  2.52 pg m<sup>-3</sup> yr<sup>-1</sup>). This indicates that they are representative of the SH 336 background.

Regionalisation with 4-days backward trajectories provide only a few monthly values for "Local" Because less than 1% of measurements could be attributed to this class. The trends for the remaining regions provide a similar pattern as regionalisation with 3 or 5 days backward trajectories. The highest and lowest  $CO_2$  trends were found for "Continental" and "South American", respectively, both when calculated from monthly averages and medians. The highest  $CH_4$  trends were observed for "Continental" when calculated from monthly averages and for "Eastern Ocean" when calculated from monthly medians. The second highest  $CH_4$ trends were for "Eastern Ocean" in the former and "Continental" in the latter case. The lowest and second lowest  $CH_4$  trends were found for "Antarctic" and "South American" air masses. Similarly to 3-days regionalisation, the GEM trends for "Antarctic" and "South American" were comparable and significantly larger than those for "Atlantic". Only the difference between the former two and the latter one is smaller than in the 3-days regionalisation.





349 In summary, the patterns of GEM,  $CO_2$ , and  $CH_4$  trend differences provide another piece of 350 evidence for an oceanic GEM source.

### 351 3.4 Regional abundance

352 In order to determine the impact of regional source and sink regions on mercury 353 concentrations at Cape Point and to evaluate whether observations at this location are 354 representative for the southern hemisphere in general, we investigate the abundance of air 355 masses from different source and sink regions and their impact on observed average 356 concentrations and long-term trends. Air masses from long range transport (Atlantic, Antarctic, 357 South American) make up 90% of all air masses observed at Cape Point. Seasonally averaged 358 observed concentrations from these regions show a high correlation with the averages of all 359 observations at Cape Point with R<sup>2</sup> values mostly above 0.9 (Table 2). Only Antarctic air 360 masses during austral summer and autumn exhibit a lower correlation.

Air masses from the sectors eastern ocean and continental on the other hand show very low correlations with the averages observed at Cape Point indicating that these air masses differ significantly from the rest. On average, transport from these two regions make up 10% of the air masses at Cape Point (Table 1). Their prevalence varies mostly only by 1 to 2 percentage points from year to year with a peak of 10% continental air masses in 2011. However, we found that the prevalence of air masses from source and sink regions is not the driver of the inter-annual variability of Hg concentrations at Cape Point. (e.g. even with twice as much as average air masses from the sink region 2011 was no year with particularly low Hg concentrations) (Figs. 7, 8). Because of this and based on the comparison with measurements at Amsterdam Island (see accompanying paper) we are confident that mercury concentrations observed at Cape Point are representative for the southern hemisphere background. Additionally, based on the presented work we are able to filter out the source and sink regions from the dataset for further analysis. Figure 9 depicts the whole GEM dataset with values from source and sink regions highlighted.

	Annua I	Spring (SON)	Summer (DJF)	Autumn (MAM)	Winter (JJA)
Antarctic	0.89	0.95	0.75	0.72	0.96
South American	0.95	0.95	0.91	0.95	0.97
Continental	0.39	0.54	0.05	0.59	0.33
Eastern Ocean	0.81	0.77	0.58	0.14	0.84
Atlantic	0.98	0.97	0.90	0.94	0.99

375 Table 2: Correlation coefficient (R<sup>2</sup>) of regional average concentrations with averages of all

376 measurements at Cape Point. Values are based on monthly averages (N=30). Antarctic, Atlantic, and 377 South American air masses exhibit a high correlation with the overall mean concentrations observed at 378 Cape Point.







380 Figure 8: Annual average concentrations at Cape Point from 2007 to 2017 (black line) and regional 381 averages (colored lines). It can be seen that the minimum in 2009 and the maximum in 2014 is present 382 in all source regions.



383 Fig. 9 Complete Cape Point dataset (grey) with observations originating from source (red) and sink 384 (blue) regions superimposed. The colored x-axis parallel indicates the long-term average (black: 385 complete dataset, red: source region, blue: sink region).

### 386 3.5 Inter-annual variability

387 So far the trend in GEM concentrations observed at Cape Point and the fact that it seemingly 388 changed from increasing to stable between 2006 and 2017 is still unexplained. Having shown 389 that the observations at Cape Point are not dominated by regional processes the question 390 arises which large-scale processes modulate the signal on annual and decadal time scales.

391 At this point, our null hypothesis is that mercury concentrations in the southern hemisphere 392 were stable over the last decade but processes on global and hemispheric scales superimpose 393 a (multi-)annual modulation on the signal. Based on our analysis so far we can exclude 394 changes in atmospheric transport patterns as the cause for the inter-annual variability and 395 correct the data for local source and sink regions. Thus, in our opinion only global source 396 processes remain as possible explanations for the observed anomaly. The identified processes





397 are: marine emissions, biomass burning and gold mining which are the major sources for 398 mercury in the Southern Hemisphere, as well as volcanic activity.

Sign Especially, the low mercury concentrations observed in 2009 and the high values observed in 2014 seem to be at least partially a large-scale phenomenon. A screening of international 2014 observation networks showed that also Mace Head – which is located in the Atlantic Ocean in 2015 the northern hemisphere – also has the lowest annual average mercury concentrations in 2009 2013 and the highest in 2014. For the year 2009 the mercury emission inventory of Streets et al. 2014 (2019) postulates a sudden plummet in global gold mining activity. Comparing the annual 2015 anomaly from the ten-year average, gold mining activity are correlated with observed GEM 2016 concentrations (R=0.64). Similarly, we found a correlation with biomass burning in the 2017 Southern Hemisphere (mostly Africa) (R=0.75) (Jiang et al., 2017). We removed air masses 2018 from the identified source and sink regions from the data set and used a regression analysis to 2019 correct for changes in global gold mining and biomass burning emissions (Figure 10). The 2010 resulting signal becomes relatively flat with only two peaks remaining in 2012 and 2014. We 2011 hypothesize that volcanic emissions which were above average in 2011 and 2014 might be 2012 responsible for higher mercury concentrations in those years.



414 Fig. 10: Annual anomaly from ten-year average mercury concentrations. Original dataset at Cape Point 415 (purple), corrected for biomass burning emissions (green) (Jiang et al., 2017) and additionally corrected 416 for gold mining emissions (orange) (Streets et al., 2019).

### 417 4. Conclusions

418 Our goal was to improve the understanding of mercury cycling in the Southern Hemisphere. 419 For this, we combined ten years of GEM observations at Cape Point, South Africa with hourly 420 backward trajectories calculated with the HYSPLIT model. Our findings are that:

421 (1) The continent is the major sink region for mercury with the exception of significant point 422 sources, mostly linked to coal combustion.

423 (2) The warm Agulhas current to the south-east is the major source of atmospheric mercury 424 observed at Cape Point.





425 (3) Separating the ground-based observations into air parcels from different source regions 426 showed that mercury behaves opposite to known pollutants of anthropogenic and continental 427 origin.

428 (4) Mercury concentration in air masses from Antarctic, Atlantic, and south American origin 429 were statistically indistinguishable. We interpret these observations as a good representation 430 of the southern hemispheric background.

431 (5) We find that the trends in GEM concentrations postulated in the past are probably an 432 artifact of single years with unusually high or low GEM concentrations (see accompanying 433 paper, this SI). We were able to show that these exceptional years could be explained by 434 changes in global emissions from biomass burning and gold mining, two major sources of 435 mercury in the Southern Hemisphere.

436 (6) With the Ocean as the main source of mercury in the southern hemisphere it can be 437 expected that an increased air-sea flux due to the larger concentration gradients will 438 compensate reductions in global atmospheric emissions in the short term future.

### 439 References

440 Amos, H. M., D. J. Jacob, D. G. Streets, and E. M. Sunderland (2013), Legacy impacts of all-441 time anthropogenic emissions on the global mercury cycle, Global Biogeochem. 442 Cycles, 27, 410-421, doi:10.1002/gbc.20040. 443 Baker, P.G.L., Brunke, E.-G., Slemr, F., Crouch, A. (2002), Atmospheric mercury 444 measurements at Cape Point, South Africa. Atmos. Envron. 36 (14) 2459-2465. 445 Belelie, M.D., Piketh, S.J., Burger, R.P., Venter, A.D., Naidoo, M., 2018, Characterisation of 446 ambient Total Gaseous Mercury concentrations over the South African Highveld. 447 Atmospheric Pollut. Res. https://doi.org/10.1016/j.apr.2018.06.001 448 Brunke, E.-G., Labuschagne, C., Parker, B., Scheel, H.E., Whittlestone, S., 2004. Baseline air 449 mass selection at Cape Point, South Africa: application of 222Rn and other filter criteria 450 to CO2. Atmos. Environ. 38, 5693-5702. 451 https://doi.org/10.1016/j.atmosenv.2004.04.024 452 Brunke, E.-G., Labuschagne, C., Ebinghaus, R., Kock, H. H., Slemr, F. (2010), Gaseous 453 elemental mercury depletion events observed at Cape Point during 2007-2008. Atmos. 454 Chem. Phys., 10, 1121-1131, 2010. 455 Buishand, T.A., 1982. Some methods for testing the homogeneity of rainfall records. J. Hydrol. 58, 11-27. https://doi.org/10.1016/0022-1694(82)90066-X 456 457 Carslaw, D.C., and Ropkins, K.: openair – An R package for air guality analysis, Environ. Modelling Software, 27-28, 52-61, 2012. 458 459 Fischer, H., Pozzer, A., Schmitt, T., Jöckel, P., Klippel, T., Taraborrelli, D., and Leleieveld, J.: 460 Hydrogen peroxide in the marine boundary layer over the South Atlantic during the 461 OOMPH cruise in March 2007, Atmos. Chem. Phys., 15, 6971-6980, 2015. 462 Gschwend, P.M., Macfarlane, J., K., Newman, K.A., 1985. Volatile halogenated organic compounds released to seawater from temperate marine macroalgae. Science 227, 463 464 1033-1035. 465 Jiang, Z., Worden, J.R., Worden, H., Deeler, M., Jones, D.B.A., Arellano, A.F., and Henze, 466 D.K.: A 15-year record of CO emissions constrained by MOPITT CO observations, 467 Atmos. Chem. Phys., 17, 4565-4583, 2017.





468 Kaiser R. and Gottschalk G., (1972), Elementare Tests zur Beurteilung von Messdaten. 469 Bibliographisches Institut. 470 Kalnay, E., Kanamitsu, M., Kistler, R., Collins, W., Deaven, D., Gandin, L., Iredell, M., Saha, S., 471 White, G., Woollen, J., Zhu, Y., Chelliah, M., Ebisuzaki, W., Higgins, W., Janowiak, J., 472 Mo, K.C., Ropelewski, C., Wang, J., Leetmaa, A., Reynolds, R., Jenne, R., and Joseph, 473 D.: The NCEP/NCAR 40-year reanalysis project, Bull. Amer. Meteor. Soc., 77, 437-470, 474 1996. 475 Martin, L.G., Labuschagne, C., Brunke, E.-G., Weigelt, A., Ebinghaus, R., Slemr, F., 2017. 476 Trend of atmospheric mercury concentrations at Cape Point for 1995–2004 and since 477 2007. Atmos. Chem. Phys., 17, 2393-2399. 478 Monks, P.S., Carpenter, L.J., Penkett, S.A., Ayers, G.P., Gillett, R.W., Galbally, I.E., and Meyer, 479 C.P.: Fundamental ozone photochemistry in the remote marine boundary layer: the 480 SOAPEX experiment, measurement and theory, Atmos. Environ., 32, 3647-3664, 1998. 481 NOAA Air Resources Laboratory (ARL), 2004, http://ready.arl.noaa.gov/gdas1.php, Tech. rep. 482 Slemr, F., Brunke, E.-G., Labuschagne, C., and Ebinghaus, R.: Total gaseous mercury concentrations at the Cape Point GAW station and their seasonality, Geophys. Res. 483 484 Lett., 35, L11807, doi:10.1029/2008GL033741, 2008. 485 Slemr, F., Brunke, E.-G., Whittlestone, S., Zahorowski, W., Ebinghaus, R., Kock, H.H., and Labuschagne, C.: 222 Rn-calibrated mercury fluxes from terrestrial surface of southern 486 487 Africa, Atmos. Chem. Phys., 13, 6421-6428, 2013. 488 Slemr, F., Weigelt, A., Ebinghaus, R., Bieser, J., Brenninkmeijer, C.A., Rauthe-Schöch, A., 489 Hermann, M., Martinsson, B.G., van Velthoven, P., Bönisch, H., Neumeier, M., Zahn, A., 490 Ziereis, H., 2018. Mercury distribution in the upper troposphere and lowermost 491 stratosphere according to measurements by the IAGOS-CARIBIC observatory: 2014-492 2016. Atmos. Chem. Phys., 18, 12329-12343 493 Stein, A. F.; Draxler, R. R.; Rolph, G. D.; Stunder, B. J. B.; Cohen, M. D.; Ngan, F. NOAA's 494 HYSPLIT Atmospheric Transport and Dispersion Modeling System. Bull. Amer. Meteor. 495 Soc. 2015, 96 (12), 2059-2077. 496 Streets, D.G., Horowitz, H.M., Lu, Z., Levin, L., Thackray, C.P., and Sunderland, E.M.: Global 497 and regional trends in mercury emissions and concentrations, 2010- 2015, Atmos. 498 Environ., 201, 417-427, 2019. 499 UNEP: The Minamata Convention on Mercury, available at: 500 http://www.mercuryconvention.org/Countries/tabid/3428/language/en-US/Default.aspx, 501 (last access: June 2019), 2013. 502 Venter, A.D., Beukes, J.P., van Zyl, P.G., Brunke, E.-G., Labuschagne, C., Slemr, F., 503 Ebinghaus, R., Kock, H., 2015. Statistical exploration of gaseous elemental mercury 504 (GEM) measured at Cape Point from 2007 to 2011. Atmos Chem Phys 15, 10271-505 10280. https://doi.org/10.5194/acp-15-10271-2015 506 Weigelt, A., Ebinghaus, R., Manning, A.J., Derwent, R.G., Simmonds, P., Spain, T.G., Jennings, S.G., Slemr, F.: Analysis and interpretation of 18 years of mercury 507 508 observations since 1996 at Mace Head, Ireland, Atmospheric Environment 100, 85-93, 509 2015. 510 Zahorowski, W., Chambers, S.D., and Henderson-Sellers, A.: Ground-based radon-222 observations and their application to atmospheric studies, J. Environ. Radioactivity, 76, 511 512 3-33, 2004.