

1 **Measurement report: Long-term variations in carbon monoxide at a**
2 **background station in China's Yangtze River Delta region**

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30 **Abstract**

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32 This study analyzed the long-term variations in carbon monoxide (CO) mixing ratios from January
33 2006 to December 2017 at the Lin'an regional atmospheric background station (LAN; 30.3°N,
34 119.73°E, 138 m a.s.l.) in China's Yangtze River Delta (YRD) region. The CO mixing ratios were
35 at their highest (0.69 ± 0.08 ppm) and lowest (0.54 ± 0.06 ppm) in winter and summer, respectively.
36 The average daily variation of CO exhibited a double-peaked pattern, with peaks in the morning
37 and evening and a valley in the afternoon. A significant downward trend of -11.3 ppb/yr of CO was
38 observed from 2006 to 2017 at the LAN station, which was in accordance with the negative trends
39 of the average CO mixing ratios and total column retrieved from the satellite data (the Measurements
40 Of Pollution In The Troposphere, MOPITT) over the YRD region during the same period. The
41 average annual CO mixing ratio at the LAN station in 2017 was 0.51 ± 0.04 ppm, which was
42 significantly lower than that (0.71 ± 0.12 ppm) in 2006. The decrease in CO levels was largest in
43 autumn (-15.7 ppb/yr), followed by summer (-11.1 ppb/yr), spring (-10.8 ppb/yr), and winter (-9.7
44 ppb/yr). Moreover, the CO levels under relatively polluted conditions (the annually 95th percentiles)
45 declined even more rapidly (-22.4 ppb/yr, $r = -0.68$, $p < 0.05$) from 2006 (0.91 ppm) to 2017 (0.58
46 ppm) and the CO levels under clean conditions (the annually 5th percentiles) showed a decreasing
47 evidence but not statistically significant ($r = -0.41$, $p = 0.19$) throughout the years. The long-term
48 decline and short-term variations in the CO mixing ratios at the LAN station were mainly attributed
49 to the implementation of the anthropogenic pollution control measures in the YRD region and to the
50 events like Shanghai Expo in 2010 and Hangzhou G20 in 2016. The decreased CO level may
51 influence atmospheric chemistry over the region. The average OH reactivity of CO at the LAN
52 station is estimated to significantly drop from 4.1 ± 0.7 s⁻¹ in 2006 to 3.0 ± 0.3 s⁻¹ in 2017.

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54 **Keywords:** CO, Long-term trend, Background level, the Yangtze River Delta region

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60 **1. Introduction**

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62 Carbon monoxide (CO) is a key **player** in the atmospheric carbon cycle (Novelli et al., 1992).

63 In the troposphere, CO is one of the important air pollutants with high mixing ratios. The volume

64 mixing ratios of CO can reach an order of 10^{-6} (Khalil et al., 1999). CO is also a reactive trace gas

65 that considerably affects health, ecology, and climate, and hence recommended by the Global

66 Atmosphere Watch (GAW) of the World Meteorological Organization (WMO) for priority

67 observation. Fossil fuel combustion (mainly in the northern hemisphere), biomass combustion

68 (mostly in the southern hemisphere), and natural processes (the oxidation of organic compounds,

69 such as methane [CH₄] and isoprene) are the main sources of CO (Holloway et al., 2000; Thompson

70 et al., 1986; Novelli et al., 1998; Andreae and Merlet, 2001; Bakwin et al., 1994). The major sink

71 for CO is its reaction with OH radicals in the troposphere (Holloway et al., 2000; Thompson et al.,

72 1986; Novelli et al., 1998; WMO, 2003). The lifetime of CO in the atmosphere ranges from weeks

73 to months, which makes it an ideal tracer for atmospheric transport processes (Steinfeld and Jeffrey,

74 1998; Worden et al., 2013). Because CH₄ and CO can react with OH radicals (Thompson et al., 1992;

75 Daniel and Solomon, 1998), certain CO mixing ratios can indirectly cause a decrease in CH₄ and an

76 increase in CO₂. Therefore, CO is recognized as an important indirect greenhouse gas. Moreover,

77 CO **can be** an important precursor for the photochemical generation of ozone **in the rural areas**

78 (Demerjian et al., 1972).

79 Continuous long-term observation is a method for studying large-scale CO sources, sinks, and

80 long-distance transport. This method allows the CO balance to be determined on a regional or global

81 scale (Fang et al., 2014). In the past decades, many studies have explored the long-term change in

82 CO levels through ground-, aircraft-, or satellite-based observations (Yurganov et al., 2010; Worden

83 et al., 2013; Ahmed et al., 2015; Cohen et al., 2018; Wang et al., 2018). Most of these studies have

84 revealed downward trends for CO concentration. For example, Worden et al. (2013) reported that

85 the CO total column over China decreased by $1.6\% \pm 0.5\%/yr$ from 2002 to 2012. Ahmed et al.

86 (2015) analyzed long-term CO observations at two urban sites in Seoul and reported a downward

87 trend of CO from 2004 to 2013. Wang et al. (2018) found that from 1998 to 2014, the total column
88 amount of CO over Beijing and Moscow decreased at $1.14\% \pm 0.87\%/yr$ and $3.73\% \pm 0.39\%/yr$,
89 respectively. Cohen et al. (2018) analyzed the trends of CO in the upper troposphere from 2001 to
90 2013. In their study, almost all observed trends were negative, with the estimated slopes ranging
91 from -1.37 to -0.59 ppb/yr. The CO data recorded in the Arctic ice core indicated that the CO
92 mixing ratios in this region decreased after the 1970s (Petrenko et al., 2013).

93 Ground-based background measurements are crucial for verifying the accuracy of satellite
94 observation data, reflecting the impact of human activities on air quality and climate change, and
95 evaluating the effectiveness of pollution control measures. In China, many air pollutants have been
96 emitted in very large quantities. For example, the emission of CO was estimated to be about 171 Tg
97 in 2010 (Li et al., 2017). To fight against the air pollution, the country has implemented a series of
98 emission control measures in the recent decade. The effectiveness of these measures needs to be
99 verified by observational data, in particular the data from background sites. Long-term background
100 observations over a decade are relatively scarce in China. Reports of long-term background
101 observations of CO are very limited in the literature (Meng et al., 2009; Liu et al., 2019; Zhou et al.,
102 2004; Zhang et al., 2011) and none of them present an analysis of CO variations over a decade. The
103 Yangtze River Delta (YRD) is one of the most developed regions in China. The long-term
104 observation of atmospheric background CO allows for a scientific understanding of the CO source
105 and sink cycle in this region. In this study, we present 12-year (from 2006 to 2017) ground-based
106 observations of CO at a background station in the YRD region. We analyze the long-term CO
107 variations and their determinants in the background areas of eastern China. The results of this study
108 function as scientific evidence for evaluating the effectiveness of pollution control policies and as a
109 reference for formulating practicable air pollution management and emission control measures.

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111 **2. Monitoring site and data collection**

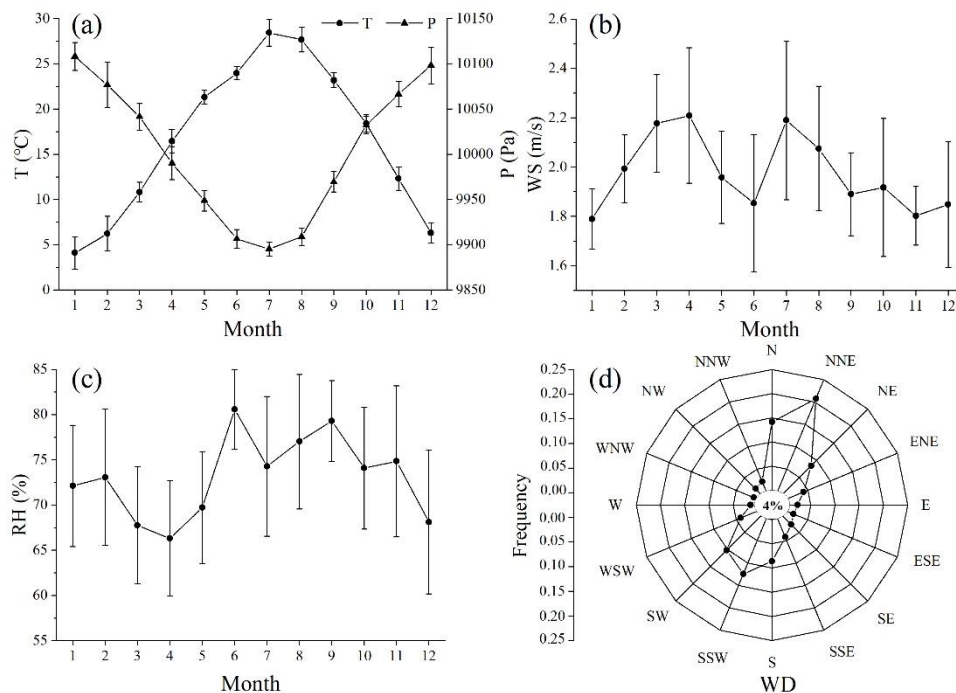
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113 The CO mixing ratios analyzed in this study were collected from January 2006 to December
114 2017 at Lin'an (LAN) station ($30^{\circ}18'$ N, $119^{\circ}44'$ E, 138.6 m a.s.l), a regional atmospheric
115 background monitoring site in China's Zhejiang province. The LAN station is one of the seven

116 atmospheric background stations operated by the China Meteorological Administration, and also a
117 member station of the World Meteorological Organization (WMO) Global Atmosphere Watch
118 (GAW) programme. The measurements at this station reflect the changes in the YRD region's
119 atmospheric background composition (Qi et al., 2012). The LAN station is located approximately
120 50 km west of Hangzhou (the capital city of Zhejiang province) and 150 km southwest of Shanghai.
121 It is influenced by a typical subtropical monsoon climate. Fig. 1 displays the seasonal variations in
122 temperature (T), air pressure (P), wind speed (WS), and relative humidity (RH) as well as the wind
123 direction (WD) frequency at the LAN station from 2006 to 2017. These data were obtained from
124 the regular meteorological observations at the LAN station. As displayed in Fig. 1, the seasonal
125 temperature trend at the LAN station was of a convex shape. The highest and lowest temperatures
126 occurred in July ($28.4 \pm 1.5^\circ\text{C}$) and January ($4.1 \pm 1.8^\circ\text{C}$), respectively. In opposition to the seasonal
127 change in temperature, the seasonal change in atmospheric pressure at the LAN station showed a
128 concave shape, with the lowest and highest pressures occurring in July ($989.51 \pm 0.77 \text{ hPa}$) and
129 January ($1010.81 \pm 1.54 \text{ hPa}$), respectively. The seasonal patterns of the WS and RH at the LAN
130 station were not as clear as those of air temperature and pressure. The seasonal average WS was
131 lowest in winter ($1.9 \pm 0.1 \text{ m/s}$) and highest in spring ($2.1 \pm 0.1 \text{ m/s}$). The RH was highest in summer
132 ($77 \pm 3\%$) and lowest in spring ($68 \pm 2\%$). The winds at the LAN station mostly originated from the
133 northeast and southwest, as shown in Fig. 1d. On average, the northeast and southwest winds
134 accounted for 29.2% and 22.6% of the winds, respectively. The calm wind frequency was 4%.

135 A gas-filter correlation infrared absorption analyzer (48C trace level, Thermo Fisher, USA)
136 was used to measure the surface CO mixing ratios. The analyzer has a limit of detection of 0.04
137 ppm. Infrared radiation is chopped and passed through a rotating gas-filter lens, half of which is
138 filled with CO and half with nitrogen. Thus, reference and measurement beams are produced in
139 alternation. The beams then pass through a narrow-band interference filter and sample cell. Because
140 the CO in the sample cell can only absorb the measurement beam, and the other gases can absorb
141 both beams, the measurement signal of CO could be obtained by comparing the attenuation intensity
142 between the reference and measurement beams.

143 The measurement signal from the CO analyzer was recorded every 5 min. Zero check and span
 144 check were conducted every 6 and 24 hours, respectively. Multipoint (>5) calibration was performed
 145 once a month using standard CO gas mixture (CO in nitrogen). Because the zero point of the
 146 instrument drifted with time, we performed linear interpolation between two adjacent zero checks
 147 to obtain the zero signals for given time point between the zero checks. These zero signals were
 148 used in the corrections of the CO data. We performed response correction according to the results
 149 of multipoint calibrations as well as the zero and span checks (Lin et al., 2009). Finally, we corrected
 150 the data according to the quantity transfer and traceability results (Lin et al., 2011). Valid 5-minute
 151 data were used to calculate the hourly mean mixing ratios. At least 10 data points were required for
 152 any given hour to calculate that hour's mixing ratio. Missing data were caused by the malfunction
 153 of the instrument from February 1 to 13, 2007, and from abnormal measurement fluctuations from
 154 May 30 to July 17, 2009.



155 **Fig. 1. Seasonal variations in (a) temperature, air pressure, (b) WS, (c) RH, and (d) WD frequency**
 156 **distribution** (the static wind frequency was 4%) at the LAN station from 2006 to 2017 (an error
 157 bar represents one standard deviation)

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 159

160 3. Results and discussion

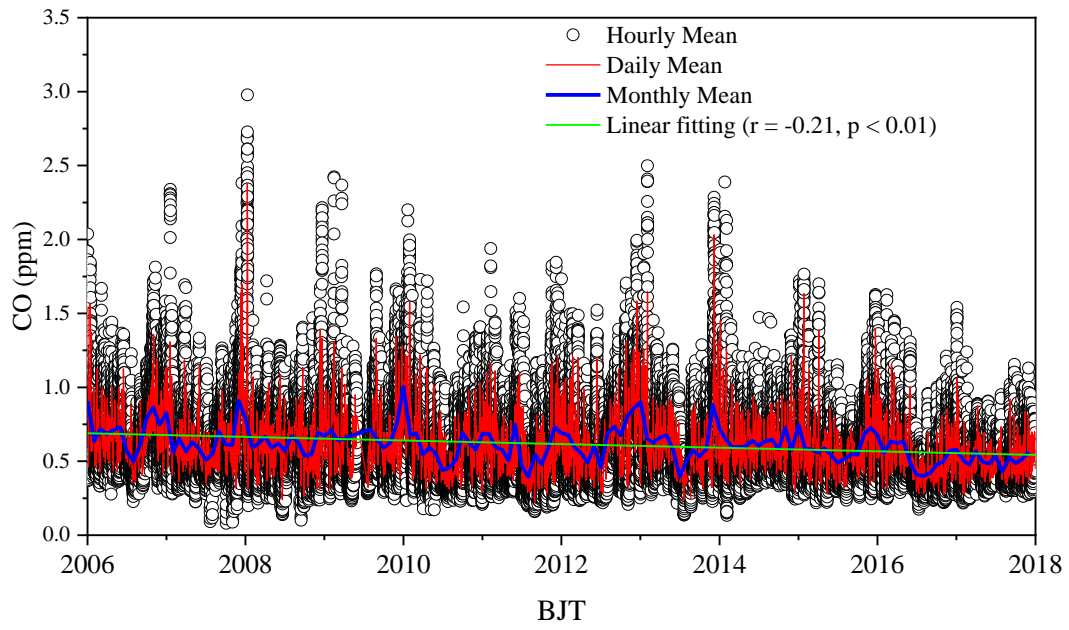
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162 3.1 Observed levels and comparisons with other sites

163 Fig. 2 displays the time series of hourly mean CO levels at the LAN station from January 1,
164 2006, to December 31, 2017 and the linear fitting results of the hourly mean CO mixing ratios. The
165 overall mean (\pm one standard deviation) and median values of the CO mixing ratios in the 12 years
166 were 0.62 (\pm 0.23) ppm and 0.57 ppm, respectively. The highest (2.98 ppm) and lowest (0.08 ppm)
167 hourly mean mixing ratios occurred at 17:00 on January 10, 2008, and 18:00 on October 4, 2007,
168 respectively. The highest hourly mean CO mixing ratio was considerably lower than the second-
169 level hourly limit (approximately 8 ppm) of the ambient air quality standard in China (GB 3095-
170 2012). The highest (2.38 ppm) and lowest (0.23 ppm) daily mean mixing ratios occurred on January
171 10, 2008, and August 31, 2011, respectively. The highest daily mean value was also below the daily
172 limit for air quality standard (3.2 ppm). The lowest monthly average CO concentration was 0.39
173 ppm on August 2011, and the highest concentration was 1.00 ppm on January 2010. The median of
174 daily mean CO levels from January 2006 to December 2017 was 0.58 ppm. The overall CO
175 concentrations at the LAN were much higher than those observed at the Waliguan global baseline
176 station from 2006-2017 and some regional background stations outside China (Table 1), indicating
177 that East China has been one of the regions with high CO levels. Table 1 also presents a comparison
178 of the seasonal average CO mixing ratios at the LAN station and other background stations in the
179 world from 2006 to 2017. The seasonal CO mixing ratios at the LAN station were marginally lower
180 than those at the Shangdianzi station in northern China (Meng et al., 2009), but were almost 3 times
181 higher than those at many other regional atmospheric background stations outside China, such as
182 the Tae-ahn Peninsula station in Korea, Yonagunijima station in Japan, Park Falls (WI) station in
183 the U.S., and Payerne station in Switzerland from 2006 to 2017 (Table 1). Moreover, the CO mixing
184 values observed at the LAN station were nearly 5 times higher than those observed at the Waliguan
185 station, a global baseline station in China. In conclusion, the CO levels at the LAN station were
186 relatively high compared to other regional atmospheric background stations outside China because
187 of more intense anthropogenic emissions in the YRD region.

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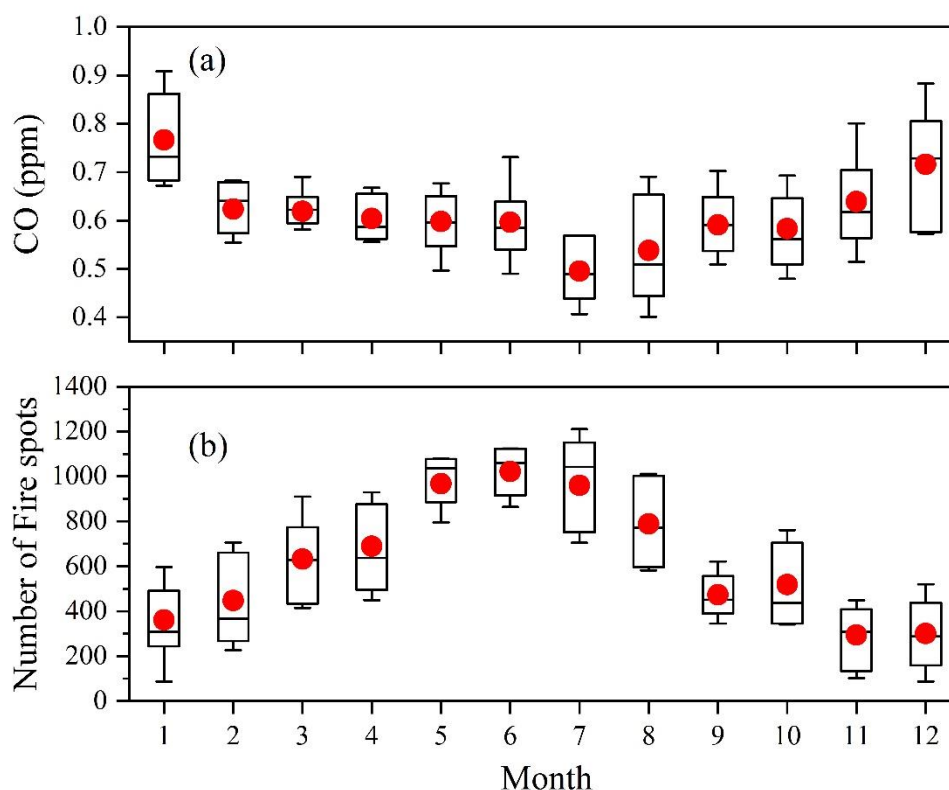
191 **Fig. 2.** Time series of the CO variations at the LAN station from 2006 to 2017

Table 1. Comparison of seasonal average CO variations at the LAN station and other similar background stations around the world

Site	Location	Period	Trends (ppb/yr)			Ref.		
			Spring (ppm)	Summer (ppm)	Autumn (ppm)		Winter (ppm)	
Lin'an, China	30°18'N, 119°44'E, 138 m	2006~2009	0.65 ± 0.04	0.59 ± 0.04	0.65 ± 0.08	0.75 ± 0.05	-11.3 This study	
		2010~2015	0.59 ± 0.04	0.54 ± 0.06	0.62 ± 0.07	0.70 ± 0.07		
		2016~2017	0.57 ± 0.08	0.46 ± 0.04	0.49 ± 0.03	0.56 ± 0.01		
Lin'an, China	30°18'N, 119°44'E, 189 m	2010/9~2012/2	0.47 ± 0.01	0.30 ± 0.01	0.41 ± 0.00	0.52 ± 0.01	-	Fang et al., 2014
Lin'an, China	30°18'N, 119°44'E, 189 m	2010/9~2017/5	0.38 ± 0.00	0.28 ± 0.00	0.37 ± 0.00	0.45 ± 0.00	-16.3	Liu et al., 2019
Shangdianzi, China	40°39'N, 117°07'E, 293 m	2006/1~2006/12	0.75 ± 0.16	0.64 ± 0.14	0.80 ± 0.12	0.76 ± 0.13	-	Meng et al., 2009
Shangdianzi, China	40°39'N, 117°07'E, 293 m	2011/12~2017/5	0.16 ± 0.00	0.18 ± 0.00	0.14 ± 0.00	0.16 ± 0.00	-1.3	Liu et al., 2019
Longfengshan, China	44° 44'N, 127° 36'E, 311 m	2006	0.21	0.20	0.27	0.38	-	Wu et al., 2008
Jinsha, China	29°38'N, 114°12'E, 750 m	2006/6~2007/7	0.44	0.39	0.66	0.60	-	(Lin et al., 2011)
Waliguan, China	36°28'N, 100°89'E, 3810 m	2006/1~2017/12	0.13 ± 0.01	0.13 ± 0.01	0.12 ± 0.01	0.12 ± 0.01	-0.67	WDCGG
Tae-ahn Peninsula, Korea	36.73°N, 126.13°E, 20 m	2006/1~2017/12	0.27 ± 0.03	0.19 ± 0.04	0.21 ± 0.03	0.23 ± 0.02	-0.43	WDCGG
Yonagunijima, Japan	24.47°N, 123.01°E, 30 m	2006/1~2017/12	0.18 ± 0.03	0.09 ± 0.01	0.13 ± 0.02	0.19 ± 0.02	-0.98	WDCGG
Park Falls (WI), the U.S.	45.93°N, 90.27°W, 868 m	2006/1~2017/12	0.17 ± 0.02	0.16 ± 0.03	0.14 ± 0.02	0.16 ± 0.02	-0.96	WDCGG
Payerne, Switzerland	46.81°N, 6.94°W, 490 m	2006/1~2017/12	0.20 ± 0.04	0.14 ± 0.01	0.20 ± 0.04	0.28 ± 0.05	-5.20	WDCGG

235 **3.2 Seasonal variation**

236 Fig. 3 shows the seasonal variations in CO mixing ratios at the LAN station and the number of
237 fire emissions (retrieved from the **Global Fire** Emissions Database version 4 described in Werf et
238 al., 2017) in the YRD region (22°N~ 40°N, 112°E~123°E) from 2006 to 2017.



239
240 **Fig. 3.** Seasonal variations in CO mixing ratios at the LAN station and the number of fire spots in
241 the YRD region from 2006 to 2017. The lines and dots in the box are the median and mean
242 concentrations, respectively, the box's lower and upper limits represent 25th and 75th percentiles
243 concentrations range, respectively, and the lower and upper whiskers correspond the 10th and 90th
244 percentiles values.

245 As can be seen in Fig. 3(a), the average CO mixing ratios were the highest in the winter (0.69
246 ± 0.08 ppm), followed by the spring (0.61 ± 0.05 ppm), autumn (0.61 ± 0.09 ppm), and summer
247 (0.54 ± 0.06 ppm). In the winter, because of the weak radiation, the photochemical consumption of
248 CO in the atmosphere decreased. Also, the atmospheric stability was high and the diffusion
249 conditions were unfavorable. Therefore, atmospheric CO accumulated easily and reached its
250 maximum concentration in the winter. **In comparison**, the photochemical reaction was strong in the

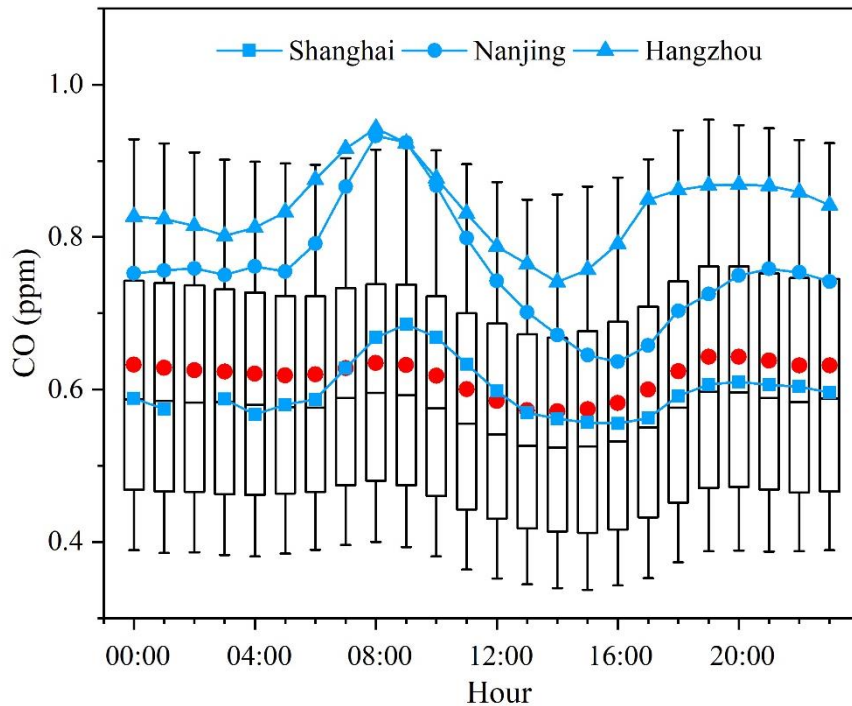
251 summer, which resulted in an increase in the mixing ratios of OH radicals and the chemical
252 consumption of atmospheric CO. Moreover, the boundary layer height was relatively high in the
253 summertime, which promoted the vertical diffusion and dilution of CO in the atmosphere. Therefore,
254 the CO mixing ratios were the lowest in the summer. By contrast, the seasonal variations in the
255 number of fire emissions in the YRD region (Fig.3b) were opposite to the trend of the CO mixing
256 ratios in different months, which indicated that open fire burning was not a main factor affecting the
257 atmospheric CO concentrations at the LAN station from 2006 to 2017.

258 **3.3 Diurnal variation**

259 The daily variations in the CO mixing ratios were influenced by emission sources, atmospheric
260 transport (horizontal and vertical), and the evolution of the atmospheric boundary layer (Xue et al.,
261 2006). Fig. 4 displays the average daily variations in the CO mixing ratios at the LAN station, along
262 with those cities Shanghai (Gao et al., 2017), Nanjing (Huang et al., 2013a) and Hangzhou (Zhang
263 et al., 2018). As displayed in Fig. 4, the CO mixing ratios exhibited double peaks, with higher CO
264 levels in the morning and evening but lower CO levels in the afternoon. The peak of the CO mixing
265 ratios at the LAN station mostly occurred in the morning (7:00–10:00) and at night (19:00–24:00).
266 The lowest CO mixing ratios were observed between 12:00 and 16:00. The hourly CO mixing ratios
267 usually reached their minimum value in the afternoon due to the high atmospheric boundary layer,
268 intense vertical diffusion mixing, and sufficient OH radicals at that time (Fang et al., 2014). The
269 Planetary Boundary Layer Height (PBLH) is a key indicator of atmospheric mixing state. As shown
270 in Fig. S1 and Fig. S2, the PBLH was rather high during the daytime and usually reached its highest
271 around 14:00, which indicated that the pollutants in the atmosphere were well mixed in the afternoon
272 and corresponded to the time when the lowest CO mixing ratios were observed (Fig. 4.). Since the
273 diurnal variations in the PBLHs at 4 sites were almost the similar according to the hourly resolution
274 (Fig. S1 and Fig. S2), the little phase shift in the CO mixing ratio peak between different sites was
275 likely attributed to the difference in local emissions. The peak CO mixing ratios at the LAN station
276 occurred during the morning and evening rush hours. This is consistent with those observed in the
277 urban areas of Shanghai (Gao et al., 2017), Nanjing (Huang et al., 2013a), and Hangzhou (Zhang et
278 al., 2018) (Fig. 4). Thus, the CO mixing ratios at the LAN station were affected by the pollutant
279 emissions related to transportation in the surroundings. However, the peak-valley difference of CO
280 at LAN was much smaller than those found in the cities, reflecting reduced impacts from direct

281 emissions on this background site.

282



283 **Fig. 4.** Average diurnal variations in CO mixing ratios from 2006 to 2015 in Shanghai, from January
284 2011 to December 2011 in Nanjing, from January 2013 to December 2013 in Hangzhou, and from
285 2006 to 2017 at the LAN station. The lines and red dots in the box are the median and mean CO
286 concentrations at the LAN station, respectively, the box's lower and upper limits represent 25th and
287 75th percentiles concentrations, respectively, and the lower and upper whiskers correspond the 10th
288 and 90th percentiles values.

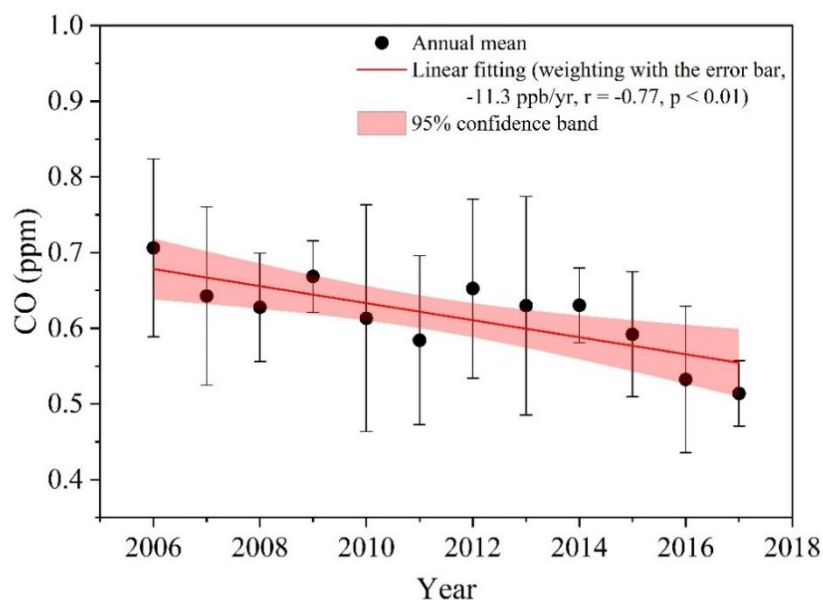
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290 3.4 Long-term trends

291 3.4.1 Trends of annual means

292 Fig. 5 shows the change in the annual mean CO mixing ratios at the LAN station from 2006 to
293 2017. The CO levels varied across the years. The World Expo was held in Shanghai from May to
294 October 2010, when air pollution prevention and control measures were strengthened in Shanghai
295 and its surrounding areas. Because of these strengthened measures, the number of days with good
296 air quality reached its highest value since 2001 (Huang et al., 2013b). Fig. 5 also indicates that the
297 average CO mixing ratio in 2010 was lower than those from 2006 to 2009 (1.5 months of data were
298 missing for the summer of 2009). The CO level continued to decline in 2011 but increased in 2012,
299 after which the CO level decreased steadily. China officially implemented the Action Plan for The

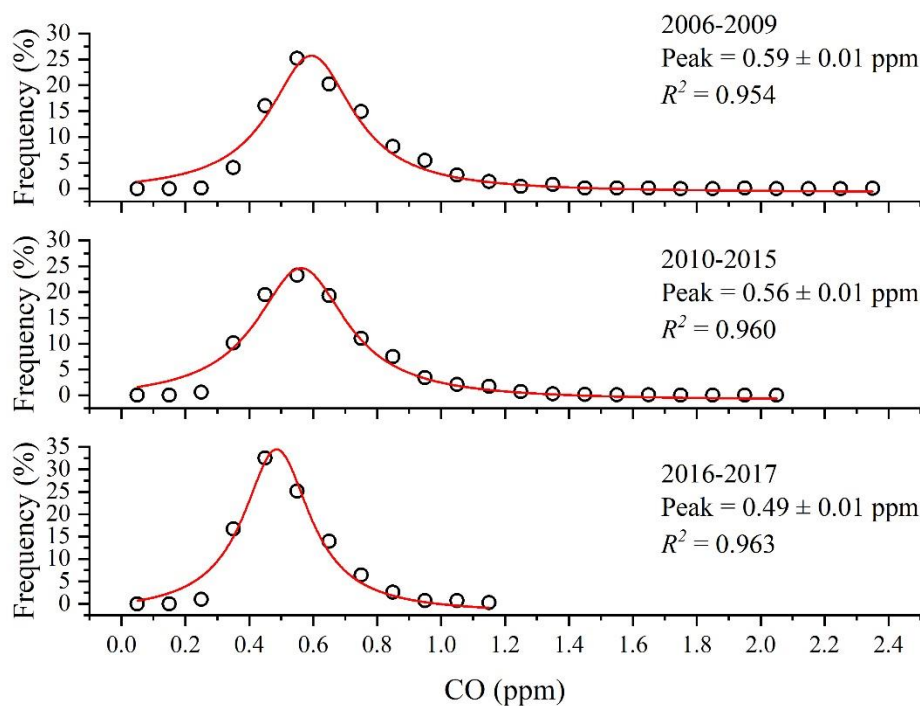
300 Prevention and Control of Air Pollution in 2013, which comprehensively intensified air pollution
 301 control efforts and reduced multi-pollutant emissions. The plan called for 5-year efforts to improve
 302 overall air quality and significantly reduce heavy pollution. As illustrated in Fig. 5, the effects of the
 303 aforementioned action plan began to be observed in 2014, and the CO mixing ratios started to
 304 decline significantly. Overall, the annual average of CO at LAN showed a decrease trend of 11.3
 305 ppb/yr ($p < 0.01$) during 2006-2017. For the period 2010-2017, we obtained a trend of -14 ppb/yr.
 306 This rate of decline in the CO mixing ratio was slightly lower than that (-16.3 ppb/yr) reported by
 307 Liu et al. (2019) for the same station for 2010-2017. The measurements of Liu et al. (2019) were
 308 performed using a cavity ring-down spectrometer, their air samples were drawn from a tower (intake
 309 height: 50 m agl), and their trend was based on non-linear fitting on CO values after removing those
 310 impacted by local events. The CO decreasing trend obtained in this study is smaller than those
 311 reported by Ahmed et al. (2015) with values of -20 ppb/yr and -13 ppb/yr respectively for two
 312 urban sites in South Korea during 2004–2013, larger than that reported by Liu et al. (2019) with a
 313 value of -1.3 ppb/yr for a regional atmospheric background station in northern China during 2011–
 314 2017, and about a factor of 2-26 of those found in regional atmospheric background stations in
 315 Korea, Japan, and Switzerland (Table 1).



316
 317 **Fig. 5.** Variation in the annual mean CO mixing ratios at the LAN station from 2006 to 2017 (the
 318 error bars represent one standard deviation calculated from monthly means)

319 Considering the variation trend in Fig. 5 and the major air pollution control policies adopted
 320 during the study period, we divided the study data into three subsets of data (collected during 2006–
 321 2009, 2010–2015, and 2016–2017, respectively). The frequency distributions of average daily CO

322 mixing ratios in the three data subsets and the Lorentz curve fitting results are displayed in Fig. 6.
 323 Approximately, a unimodal structure of CO frequency distribution was observed for all the datasets.
 324 The peak values of the Lorentz curves can be used to characterize the background concentration
 325 levels of atmospheric pollutants for a specific time and region (Lin et al., 2011). The peak of the CO
 326 Lorentz curve shifted towards lower mixing ratios over time and the trailing phenomenon of the
 327 fitting curve diminished gradually. The peak concentration of the fitting curve was 0.59 ± 0.01 ppm
 328 from 2006 to 2009. During 2010–2015 and 2016–2017, the peak CO concentrations were $0.56 \pm$
 329 0.01 and 0.49 ± 0.01 ppm, respectively. The peak frequency of the Lorentz curve was higher in
 330 2016–2017 than in 2006–2015. Moreover, the peak width was significantly narrower in 2016–2017
 331 than in 2006–2015. These are resulted from a decrease over time in the regional background mixing
 332 ratios of CO.
 333



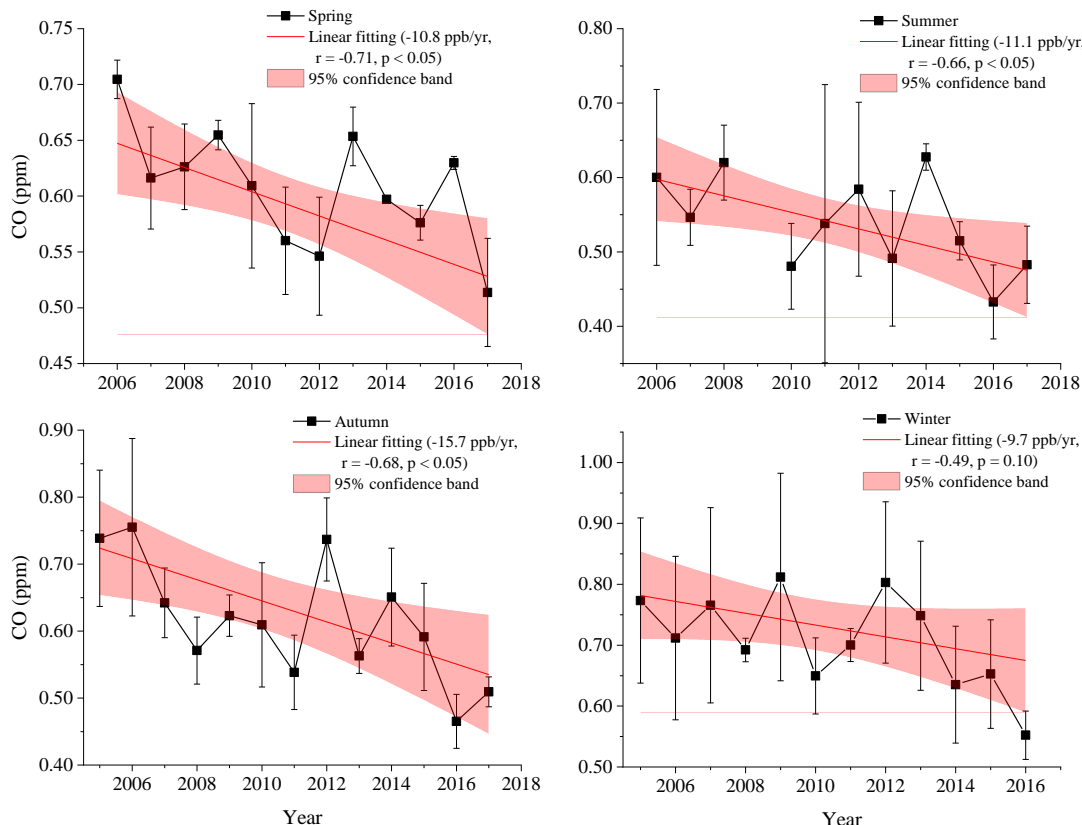
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335 **Fig. 6.** Frequency distribution of the CO mixing ratios and Lorentz curve fitting results for
 336 different time intervals

337 3.4.2 Trends of seasonal means

338 The time series of seasonal average levels of CO at the LAN station from 2006 to 2017 are
 339 displayed in Fig. 7. Linear trends were calculated from the seasonal data, with standard deviation
 340 of monthly mean values being used as weighting factors. From 2006 to 2017, the seasonal CO

341 mixing ratios exhibited larger fluctuations; nevertheless, an overall significant ($p < 0.05$) decreasing
 342 trend was observed in seasons except for the winter. The largest decrease (the slope of linear fitting)
 343 in the seasonal CO levels occurred in autumn (-15.7 ppb/yr), followed by summer (-11.1 ppb/yr),
 344 spring (-10.8 ppb/yr), and winter (-9.7 ppb/yr). As indicated in Table 1, the CO mixing ratios at the
 345 LAN station in the four seasons between 2016 and 2017 were lower than those between 2006 and
 346 2015, with the largest average decrease of 0.19 ppm occurring in winter.



347

348 **Fig. 7.** Seasonal time series and linear fitting of CO mixing ratios at the LAN station

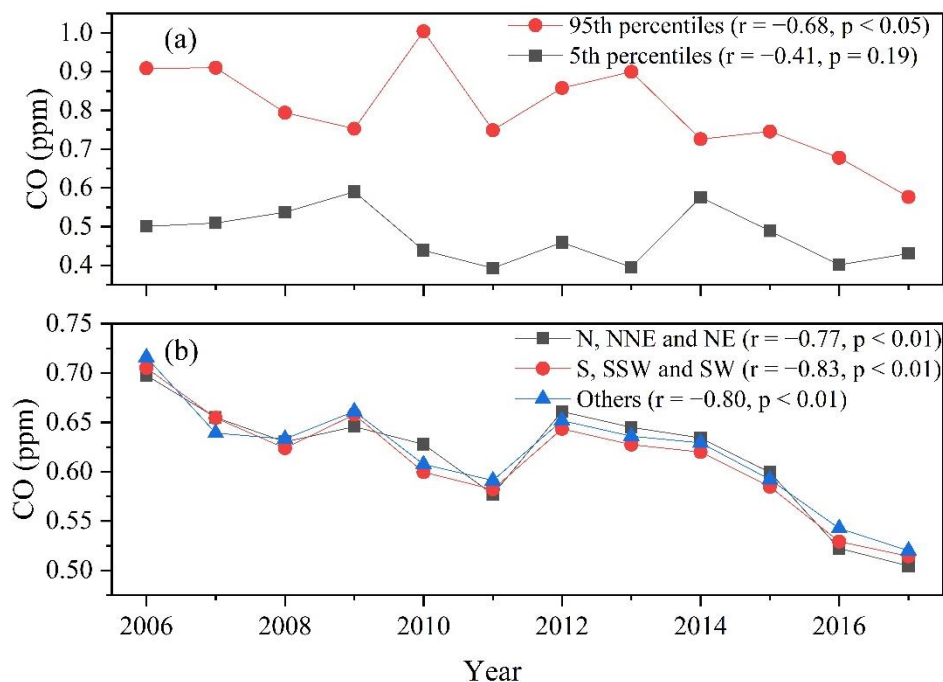
349 (Spring: March to May, Summer: June to August, Autumn: September to November, and Winter:
 350 December to February)

351

352 3.4.3 Trends of CO levels under clean and polluted condition

353 In the annual statistics, the 95th and 5th percentiles of the CO mixing ratios can be viewed as
 354 the CO levels in the most polluted and clean (background) air masses, respectively. Here, we use
 355 these two quantities to study CO trends under polluted and clean conditions, respectively, at the
 356 LAN station. As illustrated in Fig. 8 (a), the CO concentration under the polluted condition
 357 experienced a significant decreasing trend of -22.4 ppb/yr ($r = -0.68$, $p < 0.05$) from 2006 (0.91
 358 ppm) to 2017 (0.58 ppm) and that under the clean condition descended as well but not statistically

359 significant ($r = -0.41, p = 0.19$) throughout the years. This suggests that the CO levels in pollution
 360 plumes, which are highly impacted by anthropogenic emissions in the YRD region, have been
 361 reduced greatly, and the background levels of CO at the LAN station showed a decreasing evidence
 362 at the same time. Fig. 8 (b) shows the average CO concentrations from prevailing (N, NNE, NE, S,
 363 SSW and SW) and other wind directions. As can be seen in Fig. 8 (b), the annual CO levels from
 364 different wind directions generally presented similar patterns and all of them exhibited a significant
 365 ($p < 0.01$) downward trend, suggesting that the CO concentrations in the provinces and cities
 366 surrounding the LAN station have all decreased.

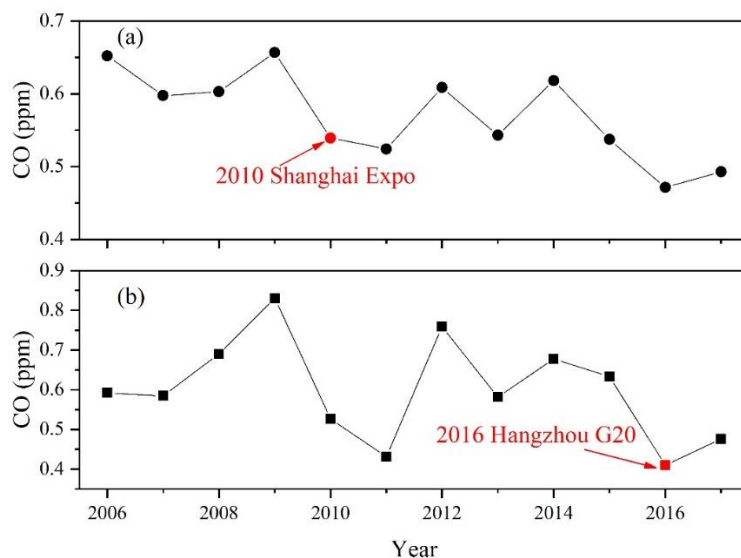


367
 368 **Fig. 8.** Trends of CO mixing ratios at 95th and 5th percentiles and from different wind directions

369
 370 **3.5 Causes and implications of the long-term variations**
 371 **3.5.1 Impacts of Shanghai Expo and G20 in Hangzhou**

372 During the Shanghai Expo in 2010 (from 1 May to 31 October) and Hangzhou G20 in 2016
 373 (from 24 July to 6 September), the Chinese government has implemented a series of joint pollution
 374 control measures in the cities of the YRD region to ensure good air quality during these mega-events.
 375 A satellite-based study (Hao et al., 2011) reported that a 12% reduction of CO concentration was
 376 observed over Shanghai city during the Expo compared to the past three years. Zhang et al. (2017)
 377 found that the ground CO levels in Hangzhou city decreased by 56% during G20 as opposed to
 378 those in 2015. In order to further evaluate the effect of these control strategies, we compared the
 379 annual trends of CO concentrations at the LAN station during the same period of Shanghai Expo

380 and Hangzhou G20, which are shown in Fig. 9 (a) and (b), respectively. The concentration of CO at
 381 the LAN station was 0.54 ppm during the Expo and 0.41 ppm during the G20, and the values were
 382 lower than those observed in Shanghai city (0.86 ppm) and Hangzhou city (0.53 ppm) in the same
 383 period. Sharp decreases (reductions of 18% during the Expo in 2010 and 35% during the G20 in
 384 2016) of the CO mixing ratios were observed at the LAN station compared to those during the same
 385 periods in the previous years. Since the meteorological conditions (the average values and standard
 386 deviations of temperature, air pressure, wind speed, relative humidity, and the wind direction
 387 frequency, see Table S1 and Fig. S3) between the during the same periods of Shanghai Expo and
 388 Hangzhou G20 and the same periods in the previous year were quite close, the results indicated that
 389 the pollution control measures worked well so as to reduce atmospheric CO concentrations in the
 390 YRD region.

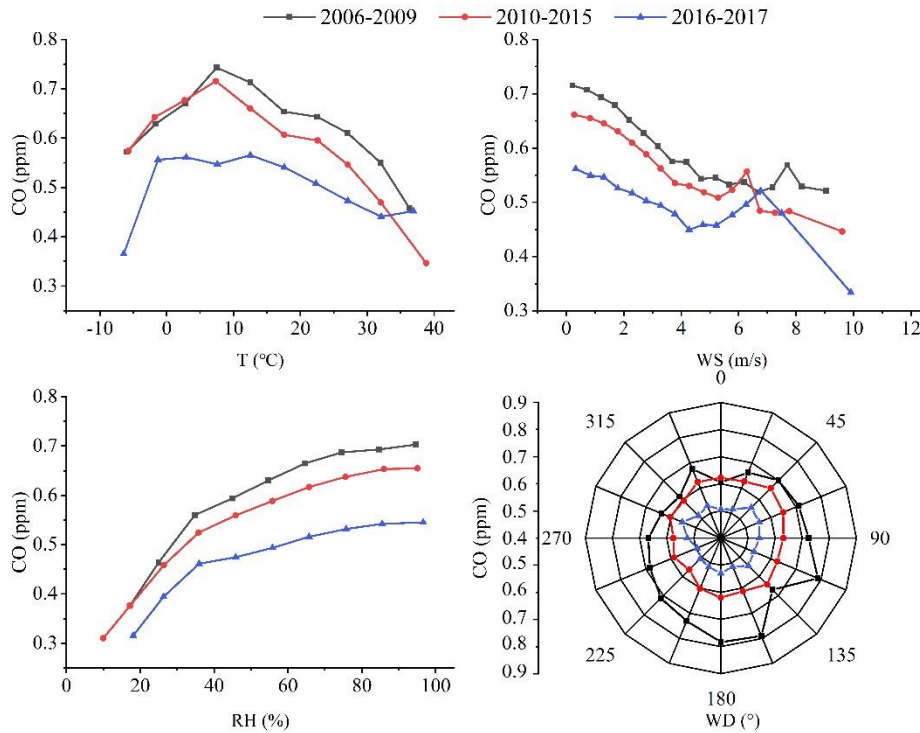


391
 392 **Fig. 9.** Average CO levels for the periods corresponding to (a) 2010 Shanghai Expo (from 1 May
 393 to 31 October) and (b) 2016 Hangzhou G20 (from 24 July to 6 September)

395 3.5.2 Relationships with meteorological conditions

396 Atmospheric CO mixing ratios are not only affected by local emission sources and the mixing
 397 ratios of atmospheric OH radicals but also by meteorological conditions. Temperature, WS, WD,
 398 and other meteorological conditions directly affect atmospheric stability and photochemical reaction
 399 intensity, which influence the diffusion, generation, consumption, and lifetime of atmospheric CO
 400 (Steinfeld and Jeffrey, 1998). Meteorological conditions varied across the years of our study period.
 401 Such variations affected the comparison of the atmospheric CO mixing ratios between different time
 402 intervals, especially when analyzing or evaluating the effectiveness of pollution control policies. To

403 minimize the effects of meteorological conditions on the analysis results, we took temperature, WS,
 404 and WD as classification variables and analyzed the variation in the CO mixing ratios under similar
 405 meteorological conditions during the three periods. The results are displayed in Fig. 10.



406
 407 **Fig. 10.** Variations of CO mixing ratios in different periods with respect to temperature (T), Wind
 408 Speed (WS), Relative Humidity (RH), and Wind Direction (WD). The intervals are 5°C, 0.5 m/s,
 409 10%, and 22.5° for T, WS, RH, and WD, respectively.

410 As displayed in Fig. 10(a), the plot of the CO mixing ratios versus the temperature showed a
 411 convex shape, with relatively low concentrations occurring at both high and low temperatures.
 412 Generally, because the photochemical reaction of CO intensifies at extremely high temperatures,
 413 and strong winds occur at extremely low temperatures, both high temperatures and strong winds
 414 can cause low CO mixing ratios. The decrease in the CO mixing ratios in a relatively high-
 415 temperature range during 2016—2017 was lower than the corresponding decreases in previous years.
 416 This result might be attributable to the summertime increase in energy consumption from the
 417 widespread use of air conditioners in China. Compared with 2006—2015, the stable area with high
 418 CO mixing ratios started to appear at lower temperatures during 2016—2017, which reflected the
 419 effectiveness of pollution control measures on the large emission sources. As displayed in Fig. 10(b),
 420 as the WS increased within a given range, the CO mixing ratios gradually decreased because of the

421 strengthened diffusion and dilution of the atmosphere. When WS increased to a given level, where
 422 this level differed between the time intervals and continually decreased overtime, the CO mixing
 423 ratios increased with WS. This may be attributable to the pollution sources being increasingly close
 424 to the LAN station because of increased urbanization over time. At a WS of 6–7 m/s, the CO mixing
 425 ratios in the different time intervals tended to be consistent. As the WS continued to increase to
 426 approximately 8 m/s, the atmospheric CO mixing ratios significantly decreased with the WS. As
 427 displayed in Fig. 10(c), the CO mixing ratios correlated positively with RH, which is consistent with
 428 the results reported by Turkoglu et al. (2004) and Ye et al. (2008). The main sink of CO is the
 429 oxidation reaction with OH radicals (Steinfeld and Jeffrey, 1998). Because water vapor is a
 430 precursor of clouds, at higher levels of RH, the atmosphere is more likely to be oversaturated with
 431 water and form clouds, and, because clouds can reflect sunlight and reduce the ultraviolet radiation
 432 reaching the ground, the photochemical reaction between CO and OH radicals is weakened (Ye., et
 433 al., 2008). Fig. 10(d) displayed the change in CO mixing ratios with respect to WD. The figure
 434 indicates that CO levels were the highest in the south sector of the LAN station.

435 Table 2 summarized the average percentage decrease in the CO mixing ratios during 2010–
 436 2015 and 2016–2017 relative to CO mixing ratios in the previous time intervals under the same
 437 meteorological conditions (temperature, WS, RH, and WD). As indicated in Fig. 10 and Table 2, the
 438 CO mixing ratios during 2016–2017 were generally lower than those during 2006–2009 and 2010–
 439 2015. Therefore, the meteorology was not the main factor contributing to the descend trend of CO.

441 **Table 2.** Comparison of the average percentage decline in CO mixing ratios during 2010–2015
 442 and 2016–2017 relative to CO mixing ratios in previous time intervals under the same
 443 meteorological factors

	Decreased Percentage (%)			
	T	WS	RH	WD
2010-2015*	-6.2	-13.6	-9.6	-11.9
2016-2017**	-14.5	-10.7	-11.7	-14.2
2016-2017*	-19.8	-16.5	-20.4	-24.4

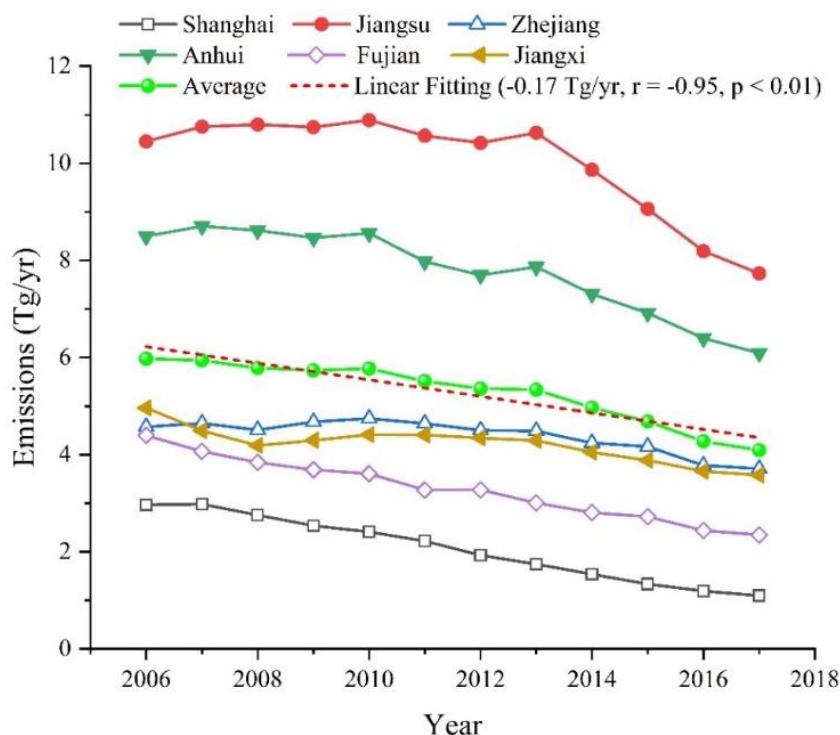
444 *: compared with 2006–2009, **: compared with 2010–2015.

445 3.5.3 Changes in emissions in neighboring provinces

446 China has implemented a comprehensive energy conservation and emission reduction policy
447 since 2006 (Zhao et al., 2008; Lei et al., 2011). Small and old factories and boilers have been
448 gradually replaced by larger and more energy-efficient alternatives. Although the focus of these
449 measures was to control sulfur dioxide emissions, these measures also greatly improved combustion
450 efficiency and thus decreased CO emissions (Zhao et al., 2012). Fig. 11 displays the change in the
451 CO emissions in six provinces and cities around the LAN station from 2006 to 2017. The emission
452 data were obtained from the Multiresolution Emission Inventory for China (Li et al., 2017). As
453 indicated in Fig. 11, the average annual CO emissions of the provinces and cities surrounding the
454 LAN station declined significantly ($r = -0.95$, $p < 0.01$), with an average decline of 170,000 tons/yr.
455 The percentages of CO emission decreased during 2016–2017 in Shanghai city as well as Jiangsu,
456 Zhejiang, Anhui, Fujian, and Jiangxi provinces were -59.3% , -25.5% , -18.6% , -27.2% , -40.1% ,
457 and -19.3% , respectively, relative to CO emission values during 2006–2009.

458 There was a strong positive correlation ($r = 0.83$, $p < 0.01$) between the annual mean CO
459 concentrations and the anthropogenic emissions of CO in the neighboring provinces. Also,
460 compared with the base year of 2006, the CO concentration in 2017 declined by 18.7%, which is
461 close to the decline value of 31.3% for the average anthropogenic emissions of CO in the
462 neighboring provinces. The decreasing percentage of the CO concentrations and the emissions were
463 overall consistent when considering larger uncertainty existing in emission. Therefore, the declined
464 trend of CO at the LAN station might be mainly attributed to the cut-down of anthropogenic
465 emissions in the YRD region.

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477 **Fig. 11.** CO emissions from 2006 to 2017 in the provinces and cities surrounding LAN station and
 478 linear fitting of the average annual CO emissions of the six provinces and cities

479 Data source: <http://meicmodel.org/dataset-mix.html>

480 3.5.4 Implications on regional atmospheric chemistry

481 The tropospheric CO has been measured on a global scale from the Measurements Of Pollution
 482 In The Troposphere (MOPITT) instrument on the spacecraft since 2000 (Deeter et al., 2017).
 483 Monthly CO mixing ratios at the surface layer and the CO total column concentrations over the
 484 YRD region from 2006 to 2017 were retrieved from MOPITT (MOP02J Version 8, 2018;
 485 http://www.satdatafresh.com/CO_MOPITT.html). We found significant correlations ($p < 0.05$)
 486 between surface CO and MOPITT CO ($r = 0.75$ and 0.61 for the MOPITT CO mixing ratio and total
 487 column, respectively) data (see Fig. S4), which indicate the good regional representativeness of
 488 Lin'an measurements. From 2006 to 2017, the average CO mixing ratio from MOPITT over the
 489 YRD region ($22.5^{\circ}\text{N} \sim 39.5^{\circ}\text{N}$, $112.5^{\circ}\text{E} \sim 123.5^{\circ}\text{E}$) in 2006 (0.11 ± 0.02 ppm) was higher than those
 490 in 2017 (0.10 ± 0.02 ppm), with a significant declining trend of -0.5 ppb/yr ($r = -0.82$, $p < 0.01$).
 491 As for the average CO total column, the value in 2006 ($1.91 \times 10^{18} \pm 0.23 \times 10^{18}$ molecules/cm²)
 492 was also higher than those in 2017 ($1.76 \times 10^{18} \pm 0.21 \times 10^{18}$ molecules/cm²), with a significant

493 declining trend of -1.07×10^{16} molecules/($\text{cm}^2 \cdot \text{yr}$) ($r = -0.70$, $p < 0.05$) from 2006 to 2017. They
494 are in consistent with the negative trends of the ground CO levels measured in the sites of the
495 WDCGG network (Table 1) and at the LAN station. Although the negative trends both in surface and
496 MOPITT CO data were found, their relative decline percent were different. Compared with the base year
497 of 2006, the surface CO declined by 1.6% annually and MOPITT CO declined by 0.4% (in mixing ratio)
498 and 0.6% (in total column), respectively.

499 The major sink for CO is reaction with OH radical (Steinfeld and Pandis, 2006), so a decrease
500 in the CO concentrations may lead to an increase in the lifetime of OH radical and thus affect the
501 atmospheric OH photochemistry (i.e., ozone production). The lifetime of OH is defined as the
502 inverse of the OH reactivity (i.e., OH loss rates), and the total OH reactivity is calculated by summing
503 over all the products of the OH reactants (CO, volatile organic compounds, nitrogen oxides, etc.)
504 concentrations times their respective rate coefficients with OH (k_{OH}) (Kovacs and Brune, 2001; Di
505 Carlo et al., 2004). The lowest average total OH reactivity ($5 \text{ s}^{-1} \sim 6 \text{ s}^{-1}$) observed in the rural areas
506 around the world (Ren et al., 2005; Ingham et al., 2009). The k_{OH} of CO is $350 /(\text{ppm} \cdot \text{min})$ at the
507 standard temperature of 298K (Vukovich, 2000) and CO generally contributed 10%~20% to the
508 total OH reactivity at the rural sites of China (Lou et al., 2010). From 2006 to 2017, the average OH
509 reactivity of CO at the LAN station exhibited a significant downward trend of $-0.07 \text{ s}^{-1}/\text{yr}$ ($r = -0.80$,
510 $p < 0.01$) and the average monthly OH reactivity of CO dropped from $4.1 \pm 0.7 \text{ s}^{-1}$ in 2006 to $3.0 \pm$
511 0.3 s^{-1} in 2017.

512

513 4. Conclusion

514 The average annual levels of CO at the LAN station during 2006–2009, 2010–2015, and 2016–
515 2017 were 0.66 ± 0.03 ppm, 0.62 ± 0.03 ppm, and 0.52 ± 0.01 ppm, respectively. From a seasonal
516 perspective, the highest seasonal average CO mixing ratio occurred in winter (0.69 ± 0.08 ppm),
517 followed by spring (0.61 ± 0.05 ppm), autumn (0.61 ± 0.09 ppm), and summer (0.54 ± 0.06 ppm).
518 The average daily variations in the CO concentration exhibited a double-peaked pattern, with high
519 CO concentrations in the morning and evening and low CO concentrations in the afternoon. Such
520 diurnal variations suggest that the CO mixing ratios at the LAN station were affected by traffic
521 pollutant emissions in its surrounding area.

522 The average annual atmospheric CO mixing ratios at the LAN station exhibited a significant
523 decreasing trend (-11.3 ppb/yr, $p < 0.01$) from 2006 to 2017, which was consistent with the negative
524 trends of the average CO mixing ratios and total column retrieved from MOPITT over the YRD
525 region. The measurements at the LAN station well reflected regional changes in atmospheric
526 background CO mixing ratios in the YRD region. The largest decrease in the CO level was observed
527 in autumn (-15.7 ppb/yr), followed by summer (-11.1 ppb/yr), spring (-10.8 ppb/yr), and winter ($-$
528 9.7 ppb/yr). The significant downward trend of the CO mixing ratios at the LAN station was not
529 caused by meteorological conditions but by strengthened pollution control measures, which
530 indicated that the adopted measures were effective. In spite of the nearly a quarter of reduction
531 during 2006-2017, the CO levels at the LAN station were still much higher than those at other
532 regional atmospheric background stations around the world so that further reductions in CO
533 emissions in the YRD region are needed. The significant decrease of regional CO level has an
534 implication for atmospheric chemistry, considering the role of CO in OH reactivity. From 2006 to
535 2017, the average OH reactivity of CO at the LAN station exhibited a significant downward trend
536 of -0.07 s⁻¹/yr ($r = -0.80$, $p < 0.01$) and dropped from 4.1 ± 0.7 s⁻¹ in 2006 to 3.0 ± 0.3 s⁻¹ in 2017.

537
538 **Data availability.** Our measurement data are deposited to an accessible repository. The data sources
539 of number of fire emissions, the annual CO emissions and the CO concentrations retrieved from
540 MOPITT over the YRD region are all listed in the reference, and the CO concentrations and the
541 meteorological data at the LAN station can be inquired about by contacting the corresponding author.

542 **Author contributions.** YJC, WLL, and BXX developed the idea for this paper and formulated the
543 research goals. QLM and JY carried out the CO field observations at the LAN station. WG provided
544 the CO data in Shanghai. YJC and WLL wrote and revised the manuscript with contributions from
545 all co-authors.

546 **Competing interests.** The authors declare that they have no conflict of interest.

547 **Acknowledgments.** This study was funded by the National Key R&D Program of China
548 (2016YFC0201900), National Natural Science Foundation of China (91744206), and Beijing
549 Science and Technology program (Z181100005418016). We thank the personnel on duty at the LAN
550 station for their assistance. This manuscript was edited by Wallace Academic Editing.

551 **References**

- 552 Ahmed, E., Kim, K. H., Jeon, E. C., and Brown, R. J. C.: Long term trends of methane, non
553 methane hydrocarbons, and carbon monoxide in urban atmosphere, *Science of the Total*
554 *Environment*, 518, 595-604, <https://doi.org/10.1016/j.scitotenv.2015.02.058>, 2015.
- 555 Andreae, M. O., and Merlet, P.: Emission of trace gases and aerosols from biomass burning,
556 *Global Biogeochemical Cycles*, 15, 955-966, <https://doi.org/10.1029/2000GB001382>,
557 2001.
- 558 Bakwin, P. S., Tans, P. P., and Novelli, P. C.: Carbon monoxide budget in the northern
559 hemisphere, *Geophysical Research Letters*, 21, 433-436,
560 <https://doi.org/10.1029/94GL00006>, 1994.
- 561 Crutzen, P. J., Heidt, L. E., Krasnec, J. P., Pollock, W. H., and Seiler, W.: Biomass burning as a
562 source of atmospheric gases CO, H₂, N₂O, NO, CH₃Cl and COS, *Nature*, 282, 253-256,
563 <https://doi.org/10.1038/282253a0>, 1979.
- 564 Cohen, Y., Petetin, H., Thouret, V., Marecal, V., Josse, B., Clark, H., Sauvage, B., Fontaine, A.,
565 Athier, G., Blot, R., Boulanger, D., Cousin, J. M., and Nedelec, P.: Climatology and long-
566 term evolution of ozone and carbon monoxide in the upper troposphere-lower stratosphere
567 (UTLS) at northern midlatitudes, as seen by IAGOS from 1995 to 2013, *Atmos. Chem.*
568 *Phys.*, 18, 5415-5453, <https://doi.org/10.5194/acp-18-5415-2018>, 2018.
- 569 Daniel, J. S., and Solomon, S.: On the climate forcing of carbon monoxide, *Journal of*
570 *Geophysical Research: Atmospheres*, 103, 13249-13260,
571 <https://doi.org/10.1029/98JD00822>, 1998.
- 572 Deeter, M., Edwards, D., Francis, G., Gille, J., Martínez-Alonso, S., Worden, H., and Sweeney,
573 C.: A climate-scale satellite record for carbon monoxide: The MOPITT Version 7 product,
574 *Atmos. Meas. Tech.*, 10, 2533-2555, <https://doi.org/10.5194/amt-10-2533-2017>, 2017.
- 575 Demerjian, K. L., Kerr, J. A., and Calvert, J. G.: The Predicted Effect of Carbon Monoxide on
576 the Ozone Levels in Photochemical Smog Systems, *Environmental Letters*, 3, 73-80,
577 <https://doi.org/10.1080/00139307209435456>, 1972.
- 578 Di Carlo, P., Brune, W. H., Martinez, M., Harder, H., Leshner, R., Ren, X., Thornberry, T., Carroll,
579 M. A., Young, V., Shepson, P. B., Riemer, D., Apel, E., and Campbell, C.: Missing OH
580 Reactivity in a Forest: Evidence for Unknown Reactive Biogenic VOCs, *Science*, 304,
581 722, <https://doi.org/10.1126/science.1094392>, 2004.
- 582 Fang, S., Zhou, L., Luan, T., Ma, Q., and Wang, H.: Distribution of CO at Lin'an Station in
583 Zhejiang Province, *Environmental Science*, 35(7), 2454-2459, 2014.
- 584 Gao, W., Tie, X., Xu, J., Huang, R., Mao, X., Zhou, G., and Chang, L.: Long-term trend of O₃
585 in a mega City (Shanghai), China: Characteristics, causes, and interactions with precursors,
586 *Science of The Total Environment*, 603-604, 425-433,
587 <https://doi.org/https://doi.org/10.1016/j.scitotenv.2017.06.099>, 2017.
- 588 Global Atmosphere Watch Programme/ The World Data Centre for Greenhouse Gases
589 (WDCGG) <https://gaw.kishou.go.jp/>
- 590 Hao, N., Valks, P., Loyola, D., Cheng, Y. F., and Zimmer, W.: Space-based measurements of air
591 quality during the World Expo 2010 in Shanghai, *Environmental Research Letters*, 6,
592 044004, <https://doi.org/10.1088/1748-9326/6/4/044004>, 2011.
- 593 Holloway, T., Levy II, H., and Kasibhatla, P.: Global distribution of carbon monoxide, *Journal*

594 of Geophysical Research: Atmospheres, 105, 12123-12147,
595 <https://doi.org/10.1029/1999JD901173>, 2000.

596 Huang, X., Huang, X., Wang, T., Zhuang, B., Li, S., Xie, M., Han, Y., Yang, X., Sun, J., Ding,
597 A., Fu, Z.: Observation and analysis of urban upper atmospheric carbon monoxide in
598 Nanjing. *China Environmental Science*, 33(9), 1577-1584, 2013.

599 Huang, Y., Wei, H., Duan, Y., Zhang, Y.: Ambient Air Quality Status and Reason Analysis of
600 Shanghai World Expo. *Environmental Monitoring in China*, 29(5), 58-63,
601 <https://doi.org/10.19316/j.issn.1002-6002.2013.05.012>, 2013.

602 Ingham, T., Goddard, A., Whalley, L. K., Furneaux, K. L., Edwards, P. M., Seal, C. P., Self, D.
603 E., Johnson, G. P., Read, K. A., Lee, J. D., and Heard, D. E.: A flow-tube based laser-
604 induced fluorescence instrument to measure OH reactivity in the troposphere, *Atmos.*
605 *Meas. Tech.*, 2, 465-477, <https://doi.org/10.5194/amt-2-465-2009>, 2009.

606 Khalil, M. A. K., and Rasmussen, R. A.: The global cycle of carbon monoxide: Trends and mass
607 balance, *Chemosphere*, 20, 227-242, [https://doi.org/https://doi.org/10.1016/0045-](https://doi.org/https://doi.org/10.1016/0045-6535(90)90098-E)
608 [6535\(90\)90098-E](https://doi.org/https://doi.org/10.1016/0045-6535(90)90098-E), 1990.

609 Kovacs, T., and Brune, W.: Total OH Loss Rate Measurement, *Journal of Atmospheric*
610 *Chemistry*, 39, 105-122, <https://doi.org/10.1023/A:1010614113786>, 2001.

611 Lei, Y., Zhang, Q., Nielsen, C., and He, K.: An inventory of primary air pollutants and CO₂
612 emissions from cement production in China, 1990–2020, *Atmospheric Environment*, 45,
613 147-154, <https://doi.org/10.1016/j.atmosenv.2010.09.034>, 2011.

614 Li, M., Zhang, Q., Kurokawa, J., Woo, J., He, K., Lu, Z., Ohara, T., Song, Y., Streets, D.G.,
615 Carmichael, G.R., Cheng, Y., Hong, C., Huo, H., Jiang, X., Kang, S., Liu, F., Su, H., and
616 Zheng, B.: MIX: a mosaic Asian anthropogenic emission inventory under the international
617 collaboration framework of the MICS-Asia and HTAP, *Atmos. Chem. Phys.*, 17, 935–963,
618 <https://doi.org/10.5194/acp-17-935-2017>, 2017.

619 Lin, H., Wang, Y., Hu, B., Zhu, R., and Wang, Y.: Observation and research on carbon monoxide
620 in the atmosphere of Beijing during the summertime of 2007, *Environmental Chemistry*,
621 28(4), 567-570, 2009.

622 Lin, W., Xu, X., Sun, J., Liu, X., and Wang, Y.: Study of the background concentrations of
623 reactive gases at Jinsha regional atmospheric background station and the impacts of long-
624 range transport. *Sci. China Earth*, 54(10), 1604-1613,
625 <https://doi.org/10.3724/SP.J.1146.2006.01085>, 2011.

626 Lin, W., Xu, X., Yu, D., Dai, X., Zhang, Z., Meng, Z., and Wang, Y.: Quality Control for
627 Reactive Gases Observation at Longfengshan Regional Atmospheric Background
628 Monitoring Station, *Meteorological Monthly*, 35(11), 93-100, 2009.

629 Lin, W., Xu, X., and Zhang, X.: The errors in the claimed concentrations of standard gases used
630 in the observation of reactive gases and recommended solutions, *Environmental Chemistry*,
631 30(6), 1-4, 2011.

632 Lin, W., Xu, X., Sun, J., Liu, X., and Wang, Y.: Background concentrations of reactive gases
633 and the impacts of long-range transport at the Jinsha regional atmospheric background
634 station, *Science China Earth*, 54, 1604-1613, 2011.

635 Liu, S., Fang, S., Liang, M., Sun, W., and Feng, Z.: Temporal patterns and source regions of
636 atmospheric carbon monoxide at two background stations in China, *Atmos Res*, 220, 169-
637 180, <https://doi.org/10.1016/j.atmosres.2019.01.017>, 2019.

638 Lou, S., F. H., F. R., Lu, K., B. B., Brauers, T., Chang, C.C., H. F., Häsel, R., Kita, K., Y. K.,
639 Li, X., M. S., Zeng, L., A. W., Zhang, Y., Wang, W., and A. H.: Atmospheric OH
640 reactivities in the Pearl River Delta - China in summer 2006, Measurement and model
641 results, *Atmos. Chem. Phys.*, 10, 11243-11260, [https://doi.org/10.5194/acp-10-11243-](https://doi.org/10.5194/acp-10-11243-2010)
642 2010, 2010.

643 Martinerie, P., Brasseur, G. P., and Granier, C.: The chemical composition of ancient
644 atmospheres: A model study constrained by ice core data, *Journal of Geophysical Research:*
645 *Atmospheres*, 100, 14291-14304, <https://doi.org/10.1029/95JD00826>, 1995.

646 Meng, Z., Ding, G., Tang, J., and Wang, S.: Characteristics of Trace Gases at Shangdianzi
647 Background Station in Autumn and Winter in Beijing, *Meteorological Science and*
648 *Technology*, 35(4), 550-557, <https://doi.org/10.3969/j.issn.1671-6345.2007.04.021>, 2007.

649 Meng, Z. Y., Xu, X. B., Yan, P., Ding, G. A., Tang, J., Lin, W. L., Xu, X. D., and Wang, S. F.:
650 Characteristics of trace gaseous pollutants at a regional background station in Northern
651 China, *Atmos. Chem. Phys.*, 9, 927-936, <https://doi.org/10.5194/acp-9-927-2009> 2009.

652 MOPITT (Measurements of Pollution in the Troposphere) Version 8 Product User's Guide,
653 https://www2.acom.ucar.edu/sites/default/files/mopitt/v8_users_guide_201812.pdf?_ga=2.33
654 903080.811721902.1572782618-1653752621.154823800.

655 Novelli, P., Steele, L., and Tans, P.: Mixing Ratios of Carbon Monoxide in the Troposphere,
656 *J. Geophys. Res.*, 97, 20731-20750, <https://doi.org/10.1029/92JD02010>, 1992.

657 Novelli, P., Masarie, K. A., and Lang, P.: Distributions and recent changes of carbon monoxide
658 in the lower troposphere, *Journal of Geophysical Research*, 103,
659 <https://doi.org/10.1029/98JD01366>, 1998.

660 Türkoğlu, N., Çiçek, İ., and Gurgun, G.: Analysis of effects of meteorological factors on air
661 pollutant concentrations in Ankara, Turkey, *Il Nuovo Cimento C*, 27,
662 <https://doi.org/10.1393/ncc/i2004-10032-0>, 2004.

663 Petrenko, V. V., Martinerie, P., Novelli, P., Etheridge, D. M., Levin, I., Wang, Z., Blunier, T.,
664 Chappellaz, J., Kaiser, J., Lang, P., Steele, L. P., Hammer, S., Mak, J., Langenfelds, R. L.,
665 Schwander, J., Severinghaus, J. P., Witrant, E., Petron, G., Battle, M. O., Forster, G.,
666 Sturges, W. T., Lamarque, J. F., Steffen, K., and White, J. W. C.: A 60 yr record of
667 atmospheric carbon monoxide reconstructed from Greenland firn air, *Atmos. Chem. Phys.*,
668 13, 7567-7585, <https://doi.org/10.5194/acp-13-7567-2013>, 2013.

669 Petrenko, V. V., Martinerie, P., Novelli, P., Etheridge, D. M., Levin, I., Wang, Z., Blunier, T.,
670 Chappellaz, J., Kaiser, J., Lang, P., Steele, L. P., Hammer, S., Mak, J., Langenfelds, R. L.,
671 Schwander, J., Severinghaus, J. P., Witrant, E., Petron, G., Battle, M. O., Forster, G.,
672 Sturges, W. T., Lamarque, J. F., Steffen, K., and White, J. W. C.: A 60 yr record of
673 atmospheric carbon monoxide reconstructed from Greenland firn air, *Atmos. Chem. Phys.*,
674 13, 7567-7585, <https://doi.org/10.5194/acp-13-7567-2013>, 2013.

675 Qi, H., Lin, W., Xu, X., Yu, X., and Ma, Q.: Significant downward trend of SO₂ observed from
676 2005 to 2010 at a background station in the Yangtze Delta region, China, *Science China*
677 *Chemistry*, 55, 1451-1458, <https://doi.org/10.1007/s11426-012-4524-y>, 2012.

678 Ren, X., Brune, W. H., Cantrell, C. A., Edwards, G. D., Shirley, T., Metcalf, A. R., and Lasher,
679 R. L.: Hydroxyl and Peroxy Radical Chemistry in a Rural Area of Central Pennsylvania:
680 Observations and Model Comparisons, *Journal of Atmospheric Chemistry*, 52, 231-257,
681 <https://doi.org/10.1007/s10874-005-3651-7>, 2005.

682 Seinfeld, J. H., Pandis, S. N., and Noone, K. J. P. T.: Atmospheric Chemistry and Physics: From
683 Air Pollution to Climate Change, 51, 88, <https://doi.org/10.1063/1.882420>, 1998.

684 Thompson, A. M.: The Oxidizing Capacity of the Earth's Atmosphere: Probable Past and
685 Future Changes, *Science*, 256, 1157, <https://doi.org/10.1126/science.256.5060.1157>, 1992.

686 Thompson, A. M., and Cicerone, R. J.: Possible perturbations to atmospheric CO, CH₄, and
687 OH, *Journal of Geophysical Research: Atmospheres*, 91, 10853-10864,
688 <https://doi.org/10.1029/JD091iD10p10853>, 1986.

689 Vukovich, F.: Weekday/Weekend Differences in OH Reactivity with VOCs and CO in
690 Baltimore, Maryland, *Journal of the Air & Waste Management Association* (1995), 50,
691 1843-1851, <https://doi.org/10.1080/10473289.2000.10464205>, 2000.

692 Wang, P., Elansky, N. F., Timofeev, Y. M., Wang, G., Golitsyn, G. S., Makarova, M. V., Rakitin,
693 V. S., Shtabkin, Y., Skorokhod, A. I., Grechko, E. I., Fokeeva, E. V., Safronov, A. N., Ran,
694 L., and Wang, T.: Long-Term Trends of Carbon Monoxide Total Columnar Amount in
695 Urban Areas and Background Regions: Ground- and Satellite-based Spectroscopic
696 Measurements, *Adv Atmos Sci*, 35, 785-795, <https://doi.org/10.1007/s00376-017-6327-8>,
697 2018.

698 Werf, G., Randerson, J., Giglio, L., Leeuwen, T., Chen, Y., Rogers, B., Mu, M., Marle, M., Morton,
699 D., Collatz, G., Yokelson, R., and Kasibhatla, P.: Global fire emissions estimates during 1997-
700 2016. *Earth Syst. Sci. Data*, 9, 697-720, <https://doi.org/10.5194/essd-9-697-2017>, 2017.
701 data source: <https://www.geo.vu.nl/~gwerf/GFED/GFED4/>, 2017.

702 Worden, H. M., Deeter, M. N., Frankenberg, C., George, M., Nichitiu, F., Worden, J., Aben, I.,
703 Bowman, K. W., Clerbaux, C., Coheur, P. F., de Laat, A. T. J., Detweiler, R., Drummond,
704 J. R., Edwards, D. P., Gille, J. C., Hurtmans, D., Luo, M., Martínez-Alonso, S., Massie, S.,
705 Pfister, G., and Warner, J. X.: Decadal record of satellite carbon monoxide observations,
706 *Atmos Chem Phys*, 13, 837-850, <https://doi.org/10.5194/acp-13-837-2013>, 2013.

707 WMO.: World Data Center for Greenhouse Gases Data Summary. Tokyo, Japan, WMO
708 WDCGG Report, 27, 1-92, 2003.

709 Wu, Y., Xu, H., and Yu, D.: Characteristics of CO concentrations at the Longfengshan regional
710 background station, *Environmental Chemistry*, 27(6), 847-848, 2008.

711 Xue, M., Wang, Y., and Sun, Y.: Measurement on the Atmospheric CO Concentration in Beijing,
712 *Environmental Science*, 27(2), 200-206, 2006.

713 Xia, L., Zhou, L., Tans, P., Liu, L., Wang, H., Luan, T., and Zhang, G.: Atmospheric CO₂ and
714 its $\delta^{13}\text{C}$ measurements from flask sampling at Lin'an regional background station in China,
715 *Atmospheric Environment*, 117, <https://doi.org/10.1016/j.atmosenv.2015.07.008>, 2015.

716 Yan, X., Ohara, T., and Akimoto, H.: Bottom-up estimate of biomass burning in mainland China,
717 *Atmospheric Environment*, 40, 5262-5273,
718 <https://doi.org/10.1016/j.atmosenv.2006.04.040>, 2006.

719 Yang, J., Liu, J., Yang, G., and Xu, J.: Characteristics and changing trend of carbon monoxide
720 pollution in Tianjin, *Environmental science and management*, 37(6), 89-104, 2012.

721 Yang, D., Yu, X., and Li, X.: Distribution characteristics of trace gas concentrations at the Lin
722 'an background station and the effect on aerosols, *Journal of Applied Meteorology*, 6(4),
723 400-406, 1995.

724 Ye, F., An, J., Wang, Y., and Y. J.: Analysis of O₃, NO_x, CO and correlation meteorological
725 factor at the ground layer in Beijing, *Ecology and Environment*, (4), 115-122, 2008.

726 Zhang, F., Zhou, L. X., Novelli, P. C., Worthy, D. E. J., Zellweger, C., Klausen, J., Ernst, M.,
727 Steinbacher, M., Cai, Y. X., Xu, L., Fang, S. X., and Yao, B.: Evaluation of in situ
728 measurements of atmospheric carbon monoxide at Mount Waliguan, China, *Atmos. Chem.*
729 *Phys.*, 11, 5195-5206, <https://doi.org/10.5194/acp-11-5195-2011>, 2011.

730 Zhang, G., Xu, H., Qi, B., Du, R., Gui, K., Wang, H., Jiang, W., Liang, L., and Xu, W.:
731 Characterization of atmospheric trace gases and particulate matter in Hangzhou, China,
732 *Atmos Chem Phys*, 18, 1705-1728, <https://doi.org/10.5194/acp-18-1705-2018>, 2018.

733 Zhang, L., Fu, C., Zheng, X., Ye, D.: Analysis of air pollution characteristics and influencing
734 factors in urban areas of Guangzhou. *Ecology and Environment*, 16(2), 305-308, 2007.

735 Zhou, L., Wen, Y., Li, J., Tang, J., and Zhang, X.: Background variation in atmospheric carbon
736 monoxide at Mt. Waliguan, China, *Acta Scientiae Circumstantiae*, 24(4), 637-642, 2004.

737 Zhao, H., Zheng, Y., Wei, L., Guan, Q., and Wang, Z.: Evolution and evaluation of air quality
738 in Hangzhou and its surrounding area during the G20 summit. *China Environmental*
739 *Science*, 37, 2016–2024, 2017.

740 Zhao, Y., Wang, S., Duan, L., Lei, Y., Cao, P., and Hao, J.: Primary air pollutant emissions of
741 coal-fired power plants in China: Current status and future prediction, *Atmospheric*
742 *Environment*, 42, 8442-8452,
743 <https://doi.org/https://doi.org/10.1016/j.atmosenv.2008.08.021>, 2008.

744 Zhao, Y., Nielsen, C. P., McElroy, M. B., Zhang, L., and Zhang, J.: CO emissions in China:
745 Uncertainties and implications of improved energy efficiency and emission control,
746 *Atmospheric Environment*, 49, 103-113,
747 <https://doi.org/https://doi.org/10.1016/j.atmosenv.2011.12.015>, 2012.

748 Zhu, L., Wang, G., and Zhang, Y.: Spatial and temporal distribution characteristics of the crop
749 straw resources in the Yangtze river delta region. *Agricultural Science of Guizhou*, 2017(4),
750 138-142.

751 Ambient air quality standard (GB 3095-2012), Ministry of ecology and environment, PRC.
752 http://kjs.mee.gov.cn/hjbhzbz/bzwb/dqhjbh/dqhjzlbz/201203/t20120302_224165.shtml

753 Action plan for the prevention and control of air pollution of the People's Republic of China.
754 http://www.gov.cn/zhengce/content/2013-09/13/content_4561.htm