# Response to Reviewer #1

We are grateful to the reviewer for the helpful comments and guidance that have led to important improvements of the original manuscript. Our point-by-point responses are listed below. Reviewer's comments are in black font, and authors' responses are in dark blue. Page and line numbers refer to the discussion paper acp-2020-575. The revised figures and Supplementary Information are attached in the end.

This work looks at the modeled vs. observed trends in sulfate and organic aerosol in the southeast US and finds that the model OA is overly dependent on sulfate, due to the impact of sulfate on acid-catalyzed reactive uptake of epoxydiols. The authors use the month-to-month variability in OA during summers to provide a new constraint on SOA formation mechanisms. Overall, this paper is well-written and within the scope of ACP and should be published after minor revisions.

# Major comments.

1. Page 8, line 151 - The statement "The 2x2.5 simulations are adequate when modeling continental boundary layer chemistry (Yu et al., 2016)" is true for ozone and NOx, but not for isoprene oxidation pathways. Yu et al. showed a large shift from high-NO<sub>x</sub> to low-NO<sub>x</sub> oxidation (Fig. 3), which would impact IEPOX production and SOA from isoprene. Please address this limitation of your work, and how an overestimate in the high NO<sub>x</sub> pathway would impact your results. Possibly a small sensitivity study is warranted to address this issue. Could it be that the model actually would have even more IEPOX SOA at high resolution and the problem would be worse?

Figure 3 in Yu et al. actually showed a shift from low-NO<sub>x</sub> to high-NO<sub>x</sub> regime when switching from a coarse grid to a fine grid, which may help alleviate the issue of IEPOX-SOA overestimation in this paper (because IEPOX is a product from the low-NO<sub>x</sub> pathway of isoprene oxidation). Due to the memory limitation of our local supercomputer system, we were unable to conduct a nested simulation at fine resolution using the complexSOA scheme.

We remove the sentence citing Yu et al. (2016) in the model description in Section 2.2.1 and add the following discussion in Section 4 Line 460: "The CT\_H01 scheme ... may serve as a simplified representation for climate models. **Simulations in this study are conducted at a horizontal resolution of 2°×2.5°**, which is comparable to most global climate models. However, as shown by Yu et al. (2016), from coarse to fine horizontal resolution, there will be a shift from low-NO<sub>x</sub> to high-NO<sub>x</sub> pathway for isoprene oxidation. Therefore, using a fine resolution may reduce the production of IEPOX and IEPOX-SOA, which needs further investigation. For all kinds of models, long-term filter-based measurements, especially intraseasonal MMV, are important observational constraints that should be considered in model development."

2. Is there a reason you didn't use observations past 2013?

We use the 2000-2013 observations to compare with our model results because we use the NEI2011 emission inventory in our model simulations, which does not provide annual emission

scaling factors after year 2013. We may update to NEI2014 which includes information for emissions after 2013 in future studies. As the main goal of this study is to show and explain the discrepancies between model and observations in early 2000s, we think using the current NEI2011 emission inventory is sufficient for this study.

3. Could you comment on how the lack of nonvolatile cations in ISORROPIA might impact your calculated aerosol acidity and generally effect your results? (https://pdfs.semanticscholar.org/f642/9d2c07179b7624c795aed5bf37b20aa0e2a8.pdf)

Theoretically, the lack of nonvolatile cations (NVCs) in ISORRPIA may lead to a higher aerosol ammonium-to-sulfate ratio and a lower aerosol pH (i.e. more acidic), based on Guo et al., 2018. In that paper, Guo et al. did a sensitivity study by applying observation-inferred Na<sup>+</sup> in ISORROPIA in the southeast US, which improved the trend of ion ratios but did not change the trends in ISORROPIA-predicted pH using fixed Na<sup>+</sup> concentration (Figure 6b in Guo et al., 2018).

The GEOS-Chem v12.1.1 does include NVCs (e.g. Na<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup> from anthropogenic and sea salt aerosol) in ISORROPIA (Pye et al., 2009, Pye et al., 2020), but they are not validated against measurements. In our study, the high H<sup>+</sup> concentration (low aerosol pH) in August 2000-2008 is mainly caused by low NH<sub>3</sub> emissions and high sulfate. The uncertainties associated with nonvolatile cations may lead to perturbations in aerosol pH and H<sup>+</sup> concentration, which matters for the formation of IEPOX-SOA. However, as stated in Guo et al., the contribution of NVCs is relatively small (and may introduce more uncertainties), and aerosol pH is not very sensitive to NVCs, the extent to which the modeled OA would be impacted by considering NVCs might be small.

In the methods Section 2.2.1 at Line 144, we modify the sentence to: "GEOS-Chem v12.1.1 considers sulfate, nitrate, and ammonium from all sectors, and fine-mode Na<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, Cl<sup>-</sup> from anthropogenic and sea salt sources, and employs the ISORROPIA II thermodynamic model (Fountoukis and Nenes, 2007; Pye et al., 2009; Song et al., 2018) to calculate aerosol water content and aerosol acidity (Pye et al., 2020)."

We mentioned the uncertainties associated with NVCs and acidity calculation in the results Section 3.3.2 at Line 373: "Due to the high uncertainty associated with the derived NH<sub>3</sub> emission product and acidity calculation (Guo et al., 2015, 2018; Silvern et al., 2017; Song et al., 2018; Tao and Murphy, 2019), we conducted another simulation 'CT\_H01' that fix  $a_{H+}$  level at 0.1 mol/L when calculating IEPOX uptake rate ..."

In the discussion Section 4 at Line 496: "The NH<sub>3</sub> emissions (which are critical for the calculation of aerosol acidity) are highly uncertain (Dammers et al., 2019), and the acidity calculation is further complicated by non-volatile cations (Guo et al., 2018) and meteorological conditions (Guo et al., 2015; Tao and Murphy, 2019)."

Minor comments.

4. Page 1, line 34 – The statement "Biogenic SOA may account for 60-100% of OA" is confusing, if you are going to call it biogenic SOA, but then discuss anthropogenic influence,

maybe call it something else, like biogenic-derived SOA or similar. Also, Kim et al., 2015 say that biogenic isoprene and monoterpenes account for 60% of OA, where do you get the 100% value?

Biogenic SOA is a term that usually refers to SOA formed from oxidation of biogenic volatile organic compounds. Over forest-covered regions biogenic SOA may account for up to 100% (e.g. Figure 2 in Xu et al., 2015).

To avoid confusion, we change the sentence to "Biogenic SOA (formed from atmospheric oxidation of BVOCs) may account for 60-100%..."

5. Page 5, line 70 – What about the role of NOx (and nitrate) in regulating aerosol water?

Nitrate contributes to a very small fraction of fine particles in the southeast US in summer. To be more comprehensive we add "NO<sub>x</sub> plays a complex role in regulating oxidation capacity, different oxidation pathways **and aerosol water content through aerosol nitrate.**"

6. Figure S1 – it is implied that the CSN sites are in Figure S1, which only shows SEARCH and IMPROVE.

Figure S1 has been revised. CSN sites are included as black dots. Due to the discontinuity of measurement protocol and techniques, 2000-2013 averaged OA concentrations are not shown for CSN sites.

7. Page 7, line 126 – Where does the baseline model agricultural NH3 come from?

In the US region, the baseline model NH<sub>3</sub> emission (from all sectors including agriculture) come from the EPA's National Emission Inventory NEI11v1. In other regions NH<sub>3</sub> emission come from the Community Emissions Data System (CEDS) inventory.

This information is included in Section 2.2.1 Line 124-128: "The global anthropogenic (including agricultural) emissions are from the Community Emissions Data System (CEDS) inventory, with the US region replaced by the EPA's National Emission Inventory for 2011 (NEI11v1). The monthly mean anthropogenic emissions of CO, SO<sub>2</sub>, NO<sub>x</sub>, NH<sub>3</sub>, VOCs, OC and black carbon are scaled to the year 2011 using the ratio of EPA's national annual emission totals from 2000 to 2013 (Travis et al., 2016)."

8. Page 7 line 126 - Do you mean you scale the emissions from 2011 backwards to 2000 and forwards to 2013 using the annual emission totals from EPA? What do you mean by "mapped over 0.1x0.1"

Yes. The other years' emissions are scaled using ratios of EPA's annual emission total in each year relative to the emission in 2011. The emissions have a horizontal resolution of  $0.1^{\circ}x0.1^{\circ}$ . We removed the words "mapped over 0.1x0.1" to avoid confusion as this information is not necessary.

We change the sentence to: "The monthly mean anthropogenic emissions of CO, SO<sub>2</sub>, NO<sub>x</sub>, NH<sub>3</sub>, VOCs, OC and black carbon are scaled to the year 2011 using the ratio of EPA's national annual emission totals from 2000 to 2013 (Travis et al., 2016)."

9. Page 8, line 134 – The GEOS-Chem complex scheme is also described in great detail by Pai et al., 2020. "An evaluation of global organic aerosol schemes using airborne observations"

The Pai et al. 2020 has been added here as a reference.

10. Page 8, line 148. It is confusing to dive in and describe sensitivity studies without showing us the problem in Figure 1. I don't think it would be out of place to say that the default model, shown in Figure 1, overestimates the trend in organic aerosol largely due to ... and we run 4 sensitivity studies to address this discrepancy...

We add: "The default modeled OA shows a stronger decreasing trend from 2000 to 2013, and a large month-to-month variability in early 2000s, different from the observations (more details in Figure 1, 2 and Section 3.1). To address this model-observation discrepancy, we do four sets of  $2^{\circ} \times 2.5^{\circ}$  simulations: ..."

11. Page 9, line 154 – It would be helpful to tell the reader that these sensitivity simulations will be described in Sections 2.2.2 etc..

We add: "The sensitivity simulations are further explained in Section 2.2.2-2.2.3 and Section 3.3."

12. Page 9, line 163 – do you mean at "the" or "from the" SOAS2013 campaign (citation?)

We change the sentence to: "The default IEPOX-SOA mechanism in GEOS-Chem uses aerosolphase reaction rates from laboratory chamber studies with pure acidic inorganic particles (Gaston et al., 2014; Riedel et al., 2015), and a representative effective Henry's law constant obtained by matching the model to the observations **from the** SOAS2013 campaign (**Marais et al., 2016**), to estimate the reactive uptake coefficient  $\gamma_{IEPOX}$ ."

13. Do the authors expect that not including non-volatile cations in ISOROPPIA could impact their findings?

See response to comment #3.

14. Page 10, line 178 – Can you provide some evidence/citations for us that there might be such a coating effect? Or is the ethylene glycol supposed to be representative enough of real organic aerosol coating ammonium bisulfate? Please just clarify a little more here.

Here we add a sentence: "In the real atmosphere, inorganic aerosol is generally internally mixed with other organics. The presence of an organic coating may alter the aerosol properties and suppress the uptake of IEPOX onto acidified sulfate aerosol (Anttila et al., 2006; Gaston et al., 2014)."

The reasoning is of adding a coating effect is also explained at the beginning of Section 3.3.1 (L330-337): "The modeled IEPOX-SOA shows a much stronger sensitivity to aerosol acidity than suggested by field observations, which found weak or no correlation between observed IEPOX-SOA and derived aerosol acidity (Budisulistiorini et al., 2015; Lin et al., 2013; Worton et al., 2013; Xu et al., 2015b). Lack of consideration of organic coating effect may provide one possible explanation. In the real atmosphere, inorganic aerosol is generally internally mixed with other organics (Anttila et al., 2006; Murphy et al., 2006). The presence of an organic coating may alter the solubility and diffusion properties at the surface of inorganic particles and diminish further uptake of IEPOX."

We also add at Line 180 about the limitation of ethylene glycol and the uncertainties with the linear fitted function: "In the real atmosphere when inorganic cores are coated with more viscous SOA (Zhang et al., 2018b), coating effect may be stronger because ethylene glycol is a low viscosity material. However, this simplified linear function does not consider the decreased viscosity and reduced coating effect at higher RH conditions (which is common in summertime southeast US) (Gaston et al., 2014; Zhang et al., 2018b), and prevents further IEPOX uptake when the mass fraction of OA ( $\chi_{org}$ ) is larger than 0.7, therefore this linear function may mimic a strong coating effect even though ethylene glycol is less viscous than real atmospheric SOA. The uncertainties need to be addressed in further studies with a more realistic coating parameterization (Li et al., 2020; Schmedding et al., 2019; Zhang et al., 2019b)."

15. Page 10, line 189 - Please elaborate a little more on the findings of Jo et al., 2019 since this seems important.

Jo et al. is also a modeling study. Here this effect is because the standard GEOS-Chem assumes no organic coating, and the aerosol surface and radius only consider inorganics (sulfateammonium-nitrate). By considering organic coating, the total aerosol surface and radius both increase which tends to increase the IEPOX uptake rate constant. Jo et al. found in some conditions (high aerosol pH and high IEPOX diffusion coefficient) the IEPOX uptake rate may even increase. The coating parameterizations in Jo et al. and in our study are different.

We remove the statement "consistent with another study (Jo et al., 2019)" to avoid confusion. We modify the sentence as: "The default GEOS-Chem with no organic coating calculates surface area of inorganic aerosol (Jo et al., 2019). By adding the coating effect, the increased particle radius  $R_p$  and surface area  $S_a$  of the mixed particle will partially offset (but does not outweigh) the impact of reduced reaction probability  $\gamma_{IEPOX modified}$ ."

16. In Section 2.2.3 – How do your emissions results compare to the NH3 emissions derived from wet deposition in Paulot et al., 2014? Paulot F., Jacob, D.J., Pinder R.W., Bash J.O., Travis, K., Henze D.K., Ammonia emissions in the United States, Europe, and China derived by high-resolution inversion of ammonium wet deposition data: Interpretation with a new agricultural emissions inventory (MASAGE\_NH3), J. Geophys. Res., 119, 4,343-4,364, 2014.

This CrIS-derived product has higher NH<sub>3</sub> emissions than the estimates from Paulot et al. (2014) (see Table 1 and more details in Cao et al., 2020).

17. Page 13, line 241 – Do the sulfate trends refer to Figure 2? If so, please reference it here.

We change to: "Compared to the slow decrease in OA, a faster declining trend is found for sulfate from IMPROVE (-6.9%/year) and SEARCH (-6.7%/year) for the same period (**Figure 2**)."

18. Page 11, line 199 – Is this due to improved farming practices as output increases? Or some other explanation? I am surprised NH3 is not increasing and would assume other readers might appreciate a little more detail.

NH<sub>3</sub> emissions have no significant trend in 2000-2013. The NH<sub>3</sub> gas concentration is increasing based on other observations, but it is not what we discussed here.

To avoid confusion, we add "There is no significant trend **of NH<sub>3</sub> emissions** from 2000 to 2013 (Figure S4), consistent with other studies suggesting nearly constant NH<sub>3</sub> emissions from 2001 to 2014 (Butler et al., 2016)."

19. Section2.2.3 – It would help the reader if there was a plot of the bottom-up and top-down NH3 emissions.

We add a Figure S3 to show the map of the default NEI11 (bottom-up) and the CrIS-derived (top-down) NH<sub>3</sub> emissions in June, July and August (averaged over 2000-2013).

20. Page11, line 208 – What surface observations? Please clarify what you mean, are there separate independent NH3 surface observations you compare against? Again, a plot would be useful here. Same for the seasonal cycle. These can be in the supplement according to the author's preference.

The top-down NH<sub>3</sub> emissions are derived and validated in another paper Cao et al., 2020. Here we use the new NH<sub>3</sub> emissions as a sensitivity test to show how sensitive IEPOX-SOA is to aerosol acidity and NH<sub>3</sub> emissions. We do not include any validation plots because they might deviate from the main focus of this paper. Instead we add the citation and add the text:

"The CrIS-derived NH<sub>3</sub> emissions have been validated against surface observations of NH<sub>3</sub> concentration from the Ammonia Monitoring Network (AMoN) and NH<sub>4</sub><sup>+</sup> wet deposition measurements from the National Atmospheric Deposition Program (NADP). More details can be found in Cao et al. (2020)."

21. Page11, line217 – For clarity, please specific that these are "model" IEPOX-SOA, sulfate aerosol etc.

We change to: "In this study we did a multivariate regression analysis of **modeled** monthly IEPOX-SOA ( $\mu$ g/m<sup>3</sup>) against **modeled** sulfate aerosol ( $\mu$ g/m<sup>3</sup>), aerosol acidity  $a_{H^+}$  (mol/L) and isoprene emission (*ISOP*<sub>emis</sub> mg/m<sup>2</sup>/hr)."

22. Page 13, line 244 – Why not sample the model at the locations of the sites? Would that make a difference to your trend?

The modeled summertime OA, if sampled at the locations of sites, has an average of 6.9  $\mu$ g/m<sup>3</sup> and a trend of 5.0%/year. These results are similar to the model results averaged over the whole southeast US domain (an average of 6.7  $\mu$ g/m<sup>3</sup> and a trend of 4.9%/year). For simplicity and to be consistent with other analysis in this study, we only show the model results averaged over the whole domain here.

Here we add the text: "Modeling results are averaged over the domain  $[29^{\circ} \sim 37^{\circ}N, 74^{\circ} \sim 96^{\circ}W]$  excluding ocean grid cells (Figure S1). The 2000-2013 JJA-averaged OA from the default model is 6.7 µg/m<sup>3</sup>, higher than OA from IMPROVE and SEARCH. Modeled total OA mass decreases at a rate of 4.9%/year, about 1.9 (1.6) times faster than IMPROVE (SEARCH) OA (student's t-test p<0.001). By sampling the model results at the locations of the IMPROVE and SEARCH sites, the modeled summertime OA has an average of 6.9 µg/m<sup>3</sup> and a trend of 5.0%/year, similar to the model results averaged over the whole southeast US domain. For simplicity, we show only the domain-averaged model results in all figures and analysis."

23. Figure 1 – can you explain the periods of enhanced OA around 2006 and 2011?

The enhanced OA in 2006 and 2011 is due to high isoprene emissions (Figure 3). In Section 3.2 Line 326 we change to: "Isoprene emission contributes to the overall interannual variability, for example leading to the relatively low IEPOX-SOA in 2003-2004 and the peaks in 2000, 2006 and 2011 (Figure 3)."

24. Figure 2 – It could be helpful to include the trends on this figure similar to Figure 1.

Figure 2 has been modified to include trends of OA and sulfate.

25. Page 15, line 302 – Figure 3 is difficult to interpret, please provide a little more description of what is shown here, particularly what is mean by "divided by 1 standard deviations."

We add the following descriptions: "Figure 3 shows the standardized monthly surface IEPOX-SOA concentration, sulfate concentration, aerosol H<sup>+</sup> activity and isoprene emission from the default model. For each variable, the monthly gridded data has been first averaged over the southeast US. Then, we calculate the one standard deviation of all monthly data (June, July and August data from 2000 to 2013). Finally, the domain-averaged monthly data has been divided by its standard deviation, so the variables are standardized to be unitless and their variability can be compared directly."

26. Page 16, line 325 – Is isoprene interannual variability explained by temperature? Temperature is an important driver of aerosol partitioning, should it not be included in Table S1 and Figure 3? Does temperature explain any of the interannual variation in Figure 1?

Yes, temperature is an important factor controlling the variability of isoprene emissions and therefore SOA. In GEOS-Chem, SOA formed by aqueous-phase uptake of isoprene products

(which dominates the total OA in the southeast US as shown in Figure 1) is assumed to be nonvolatile, and temperature plays a minor role in these processes. The monoterpene-derived SOA is considered to be semi-volatile and formed by reversible gas-aerosol partitioning, which is influenced by temperature. But monoterpene-SOA contributes to a smaller fraction of OA (Figure 1). Here we focus on the variability of IEPOX-SOA (which is non-volatile), and we already consider isoprene emission as a main driver. Therefore, we do not consider temperature as an additional independent driver of IEPOX-SOA in Table S1 and Figure 3.

Earlier at Line 295 we add: "We use the GEOS-Chem model here to examine the simulated IEPOX-SOA dependence on sulfate, aerosol acidity, and emissions of isoprene which produce IEPOX at high yields under low-NO<sub>x</sub> conditions (Paulot et al., 2009). **Temperature impacts the formation of IEPOX-SOA mainly through regulating isoprene emissions but does not influence partitioning as IEPOX-SOA is treated as non-volatile in GEOS-Chem. Therefore, temperature is not examined as another driver in addition to isoprene emissions."** 

27. Page 18, line 376 – I assume that the H+ value during 2013SOAS was calculated with a thermodynamic model, although constrained by observations. Please clarify.

Yes. We change the wording here to: "we conducted another simulation 'CT\_H01' that fix  $a_{H^+}$  level at 0.1 mol/L when calculating IEPOX uptake rate, corresponding to the **predicted**  $a_{H^+}$  value (**constrained by observations**) during the 2013 SOAS campaign (Weber et al., 2016)."

28. Page 19, line 378 – Can you comment on why the model agrees much better with SOAS2013 observations and 2013 OA in Figure 1, and what the limitations are of testing your revised IEPOX uptake at the low end of the SO2 emissions trend?

Figure 1 is not about SOAS2013 campaign data. Figure 1 shows the comparison between the default model results and the long-term filter measurements of OA from surface networks IMPROVE and SEARCH. Our revised IEPOX uptake agrees well with the long-term surface measurement in 2000-2013 (Figure 4C and 4D), which is the main focus of our paper. The 2000-2013 period covers a broad range of SO<sub>2</sub> emissions. The revised IEPOX-SOA also compares well with the SOAS2013 data (Figure S6).

In Line 378, we add: "The two simulations, CT\_newNH<sub>3</sub> and CT\_H01, yield similar long-term trends of IEPOX-SOA in the southeast US (Figure S5), **and they agree better with the long-term surface OA measurements from IMPROVE and SEARCH than the default model (Figure 4C and 4D).** For the SOAS2013 campaign, the CT\_H01 scheme simulates an average IEPOX-SOA concentration of 0.74  $\mu$ g/m<sup>3</sup>, similar to 0.81  $\mu$ g/m<sup>3</sup> in the default model, and agrees well with the two independent Aerosol Mass Spectrometer measurements (0.97  $\mu$ g/m<sup>3</sup> from obs\_GT and 0.68  $\mu$ g/m<sup>3</sup> from obs\_CU, see daily time series in Figure S6)."

29. General comment – why not consider trends in ammonium? Particularly to help support your improved NH3 emissions scenario?

The long-term observation network IMPROVE does not have measurements of ammonium aerosol. SEARCH shows a downward trend of ammonium aerosol, similar to the trend of sulfate

(Silvern et al., 2017). In the model ammonium (aerosol  $NH_4^+$ ) is highly correlated with sulfate and has a very similar trend as sulfate, so we do not show the trends in ammonium here. The month-to-month variability of  $NH_3$  emissions is what matters that leads to high aerosol  $H^+$  in August in early 2000s (and therefore IEPOX-SOA). The  $NH_3$  emissions do not have significant long-term trends based on EPA's annual emission totals.

We add the trends of aerosol NH<sub>4</sub><sup>+</sup>/SO<sub>4</sub><sup>2-</sup> ratio to Figure S4, and change the text in Section 3.2 around L308: "The default NH<sub>3</sub> emissions from NEI11v1 suggest no significant long-term trend from 2000 to 2013. In general, ammonium aerosol is strongly correlated with sulfate and has a similar declining trend as sulfate (Silvern et al., 2017). However, the NH<sub>3</sub> emissions in August are about 25% lower than in June and July (Figure S4). As a result, in August before 2008, the aerosol NH<sub>4</sub><sup>+</sup>/SO<sub>4</sub><sup>2-</sup> ratio is smaller (Figure S4) and the  $a_{H^+}$  is up to 3 times higher than June, leading to high production of IEPOX-SOA in August."

30. Page 22, line 467 – By gas-aerosol partitioning, do you mean reversible uptake of isoprene oxidation products? Could you be more specific on this?

a. Generally, the final paragraph that goes through potential reasons for the remaining model discrepancies is very useful but needs additional explanation. The sentence starting "Despite a large MMV in IEPOX-SOA is confusing, please rephrase to more clearly state your meaning.

For the above questions, we change the text as follows: "Such discrepancies may suggest a more important role of SOA pathways that are less dependent on inorganic aerosol, such as **terpene-SOA formed by reversible** gas-aerosol partitioning. **Terpene-SOA is included in GEOS-Chem (yellow color in Figure 1), and contributes to 8~24% of total OA, which might be underestimated compared to recent field studies. Xu et al. (2015a) finds a large MMV in IEPOX-SOA, but the less-oxidized oxygenated OA (LO-OOA, an indicator for freshly-formed monoterpene SOA) and the more-oxidized oxygenated OA (MO-OOA, also likely from biogenic sources) have little MMV in summer months, and** they contribute to more than 50% of total OA mass in the southeast US (Xu et al., 2018). The important role of monoterpenes SOA is also confirmed by molecular level characterization of organic aerosols (Zhang et al., 2018a)."

b. Is monoterpene SOA included in the GEOS-Chem simulations? If so, is the lack of MMV captured by the model? I am generally just confused by the discussion here. Are you trying to say that other SOA pathways, not in the model, might have less MMV and therefore improve model performance of the long-term trend?

Yes, monoterpene SOA is included in GEOS-Chem (see changes above). In Figure 1 yellow color represents terpene SOA (mostly monoterpene and also some sesquiterpene, see Section 2.2.1 Line 137). Terpene SOA and anthropogenic SOA are referred to as dry SOA (Line 53). They have small MMV (Figure 2B) but they only contribute to a small fraction of total OA especially when comparing to IEPOX-SOA in early 2000s (Figure 1). We would like to strengthen that monoterpene SOA may be underestimated in GEOS-Chem, compared to the field measurements (Xu et al., 2018; Zhang et al., 2018), see Line 508-515. Other pathways that are less dependent on sulfate and aerosol acidity (and probably less MMV), not in the model now, may improve the model performance and need further examination (Line 515-525).

# **Reference:**

Anttila, T., Kiendler-Scharr, A., Tillmann, R. and Mentel, T. F.: On the reactive uptake of gaseous compounds by organic-coated aqueous aerosols: Theoretical analysis and application to the heterogeneous hydrolysis of N2O5, J. Phys. Chem. A, 110(35), 10435–10443, doi:10.1021/jp062403c, 2006.

Cao, H., Henze, D. K., Shephard, M. W., Dammers, E., Cady-Pereira, K., Alvarado, M., Lonsdale, C., Luo, G., Yu, F., Zhu, L., Danielson, C. G. and Edgerton E. S., accepted by Environ. Res. Lett., 2020.

Gaston, C. J., Riedel, T. P., Zhang, Z., Gold, A., Surratt, J. D. and Thornton, J. A.: Reactive uptake of an isoprene-derived epoxydiol to submicron aerosol particles, Environ. Sci. Technol., 48(19), 11178–11186, doi:10.1021/es5034266, 2014.

Guo, H., Nenes, A. and Weber, R. J.: The underappreciated role of nonvolatile cations in aerosol ammonium-sulfate molar ratios, Atmos. Chem. Phys., 18(23), 17307–17323, doi:10.5194/acp-18-17307-2018, 2018.

Pai, S., Heald, C., Pierce, J., Farina, S., Marais, E., Jimenez, J., Campuzano-Jost, P., Nault, B., Middlebrook, A., Coe, H., Shilling, J., Bahreini, R., Dingle, J. and Vu, K.: An evaluation of global organic aerosol schemes using airborne observations, Atmos. Chem. Phys., (20), 2637–2665, doi:10.5194/acp-2019-331, 2020.

Paulot, F., Crounse, J. D., Kjaergaard, H. G., Kürten, A., Clair, J. M. S., Seinfeld, J. H. and Wennberg, P. O.: Unexpected Epoxide Formation in the Gas-Phase Photooxidation of Isoprene, Science (80-.)., 325, 730–734, doi:10.1126/science.1174251, 2009.

Pye, H. O. T., Liao, H., Wu, S., Mickley, L. J., Jacob, D. J., Henze, D. J. and Seinfeld, J. H.: Effect of changes in climate and emissions on future sulfate-nitrate-ammonium aerosol levels in the United States, J. Geophys. Res. Atmos., 114(1), 1–18, doi:10.1029/2008JD010701, 2009.

Pye, H. O. T., Nenes, A., Alexander, B., Ault, A. P., Barth, M. C., Simon, L., Collett, J. L., Fahey, K. M., Hennigan, C. J., Herrmann, H., Kanakidou, M., Kelly, J. T., Ku, I., Mcneill, V. F., Riemer, N., Shi, G., Tilgner, A., Walker, J. T., Wang, T., Weber, R., Xing, J., Zaveri, R. A. and Zuend, A.: The Acidity of Atmospheric Particles and Clouds, Atmos. Chem. Phys. Atmos. Chem. Phys., (October), 2020.

Silvern, R. F., Jacob, D. J., Kim, P. S., Marais, E. A., Turner, J. R., Campuzano-Jost, P. and Jimenez, J. L.: Inconsistency of ammonium-sulfate aerosol ratios with thermodynamic models in the eastern US: A possible role of organic aerosol, Atmos. Chem. Phys., 17(8), 5107–5118,

doi:10.5194/acp-17-5107-2017, 2017.

Xu, L., Suresh, S., Guo, H., Weber, R. J. and Ng, N. L.: Aerosol characterization over the southeastern United States using high-resolution aerosol mass spectrometry: Spatial and seasonal variation of aerosol composition and sources with a focus on organic nitrates, Atmos. Chem. Phys., 15(13), 7307–7336, doi:10.5194/acp-15-7307-2015, 2015a.

Xu, L., Guo, H., Boyd, C. M., Klein, M., Bougiatioti, A., Cerully, K. M., Hite, J. R., Isaacman-VanWertz, G., Kreisberg, N. M., Knote, C., Olson, K., Koss, A., Goldstein, A. H., Hering, S. V., de Gouw, J., Baumann, K., Lee, S.-H., Nenes, A., Weber, R. J. and Ng, N. L.: Effects of anthropogenic emissions on aerosol formation from isoprene and monoterpenes in the southeastern United States, Proc. Natl. Acad. Sci., 112(1), 37–42, doi:10.1073/pnas.1417609112, 2015b.

Xu, L., Pye, H. O. T., He, J., Chen, Y., Murphy, B. N. and Ng, N. L.: Experimental and model estimates of the contributions from biogenic monoterpenes and sesquiterpenes to secondary organic aerosol in the southeastern United States, Atmos. Chem. Phys., 18(17), 12613–12637, doi:10.5194/acp-18-12613-2018, 2018.

Yu, K., Jacob, D. J., Fisher, J. A., Kim, P. S., Marais, E. A., Miller, C. C., Travis, K. R., Zhu, L., Yantosca, R. M., Sulprizio, M. P., Cohen, R. C., Dibb, J. E., Fried, A., Mikoviny, T., Ryerson, T. B., Wennberg, P. O. and Wisthaler, A.: Sensitivity to grid resolution in the ability of a chemical transport model to simulate observed oxidant chemistry under high-isoprene conditions, Atmos. Chem. Phys., 16(7), 4369–4378, doi:10.5194/acp-16-4369-2016, 2016.

Zhang, H., Yee, L. D., Lee, B. H., Curtis, M. P., Worton, D. R., Isaacman-VanWertz, G., Offenberg, J. H., Lewandowski, M., Kleindienst, T. E., Beaver, M. R., Holder, A. L., Lonneman, W. A., Docherty, K. S., Jaoui, M., Pye, H. O. T., Hu, W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Guo, H., Weber, R. J., de Gouw, J., Koss, A. R., Edgerton, E. S., Brune, W., Mohr, C., Lopez-Hilfiker, F. D., Lutz, A., Kreisberg, N. M., Spielman, S. R., Hering, S. V., Wilson, K. R., Thornton, J. A. and Goldstein, A. H.: Monoterpenes are the largest source of summertime organic aerosol in the southeastern United States, Proc. Natl. Acad. Sci., 201717513, doi:10.1073/pnas.1717513115, 2018a.

Zhang, Y., Chen, Y., Lambe, A. T., Olson, N. E., Lei, Z., Craig, R. L., Zhang, Z., Gold, A., Onasch, T. B., Jayne, J. T., Worsnop, D. R., Gaston, C. J., Thornton, J. A., Vizuete, W., Ault, A. P. and Surratt, J. D.: Effect of the Aerosol-Phase State on Secondary Organic Aerosol Formation from the Reactive Uptake of Isoprene-Derived Epoxydiols (IEPOX), Environ. Sci. Technol. Lett., 5(3), 167–174, doi:10.1021/acs.estlett.8b00044, 2018b.

# Response to Reviewer #2

We are grateful to the reviewer for the helpful comments and guidance that have led to important improvements of the original manuscript. Our point-by-point responses are listed below. Reviewer's comments are in black font, and authors' responses are in dark blue. Page and line numbers refer to the discussion paper acp-2020-575. The revised figures and Supplementary Information are attached in the end.

Summary: The author compared the long-term trend of organic aerosol and sulfate mass loading from field measurement with GEOS-Chem simulations in this manuscript, with the model showing a steeper decrease in the OA mass loading and larger monthto-month variability than the field data. The long-term trend of sulfate, on the other hand, was well captured by the model when comparing with the field data. By applying coating effect, constant aerosol acidity, and a different NH3 emission product, the modeling results match the field data better, suggesting further study is needed to address the weak dependence of OA formation and sulfate. The manuscript is overall clearly written, but may need to address the following aspects before publishing.

First, the author did a good job explaining the modeling results in Section 3, but it was not very clear at certain sections which modeling results agree with the field measurement and which do not. The author may need to improve clarity about the modeling-field data comparisons when describing the modeling results, especially in line 251 when the author discusses that the contribution of IEPOX-SOA to total OA mass decreases from 61% to 28% from the early 2000s to 2013. Does this modeling result agree with previous field measurements (such as the results from Xu. et al. 2015 and Budisulistiorini, et al. 2016)? Maybe the author can include a sentence or two to compare the modeling data with the field data.

Section 3.1 mainly compares modeling results with long-term surface filter measurements from IMPRAOVE and SEARCH networks, focusing on the long-term trend and month-to-month variability of OA (Figure 1 and 2). We also compare the modeling results with the SOAS2013 field campaign data later in Section 3.3.2 Line 379-385 and Figure S6.

Here we add a sentence about the contribution of IEPOX-SOA: "The contribution of IEPOX-SOA to total OA mass decreases from 61% in the early 2000s to 28% in 2013. The simulated IEPOX-SOA in 2013 compares well with previous field studies which suggested that IEPOX-SOA contributed to 18~40% in southeast US sites in summer 2013 (Budisulistiorini et al., 2016; Xu et al., 2015)."

The manuscript also concluded that coating can improve the modeling result because thinner coating may enhance the formation of IEPOX-SOA. There are a few papers that also measured/modelled the effects of SOA coating on the formation of IEPOX-SOA. For instance, Zhang et al. 2018, Jo et al., 2019, and the subsequent study by Schmedding et al. 2019 discussed the effects of pre-existing coating on the formation of IEPOX-SOA. Does the result in this manuscript using ethylene glycol agree with previous studies using authentic SOA? In addition, Jo et al. 2019 shows that the uptake of IEPOX would increase with increasing coating for most of the situations using GEOS-Chem due to increasing surface tension, contrary to a decreasing

uptake with coating effect in this manuscript. Could the author explain why the trends are different in these two studies?

Compared to the default GEOS-Chem with no coating effect at all, adding a coating effect reduces (not enhances) the production of IEPOX-SOA in the summertime southeast US (Figure 4C, orange line as compared to magenta line). Jo et al. applied the parameterizations from Zhang et al., 2018 (using monoterpene-SOA as a coating material) and is different from what we used here. In Gaston et al. (2014) which we applied in this study, ethylene glycol is a relatively low viscosity material, but the simplified linear function fitted using RH=50% conditions may mimic a strong coating effect, because it does not consider the reduced viscosity and weaker coating effect at higher RH conditions, and it assumes no IEPOX uptake when the fraction of organics is higher than 70%. As a result, the added aerosol surface and particle radius does not overweigh the impact of reduced uptake coefficient when considering coating.

In Section 2.2.2 Line 180 we add: "In the real atmosphere when inorganic cores are coated with more viscous SOA (Zhang et al., 2018b), coating effect may be stronger because ethylene glycol is a low viscosity material. However, this simplified linear function does not consider the decreased viscosity and reduced coating effect at higher RH conditions (which is common in summertime southeast US) (Gaston et al., 2014; Zhang et al., 2018b), and prevents further IEPOX uptake when the mass fraction of OA ( $\chi_{org}$ ) is larger than 0.7, therefore this linear function may mimic a strong coating effect even though ethylene glycol is less viscous than real atmospheric SOA. The uncertainties need to be addressed in further studies with a more realistic coating parameterization (Li et al., 2020; Schmedding et al., 2019; Zhang et al., 2019b). We assume all OA is coated outside the inorganic aerosol core when calculating the IEPOX reactive uptake. The default GEOS-Chem with no organic coating calculates surface area of inorganic aerosol. By adding the coating effect, the increased particle radius  $R_p$  and surface area  $S_a$  of the mixed particle will partially offset (but does not overweigh) the impact of reduced reaction probability  $\gamma_{IEPOX\_modified}$ ."

The manuscript used a fixed acidity to reduce the modelled month-to-month variation of IEPOX-SOA so that the results match better with the field data. It seems this large month-to-month variation in modeling only appeared before 2008, while the month-to-month variation decreased significantly after 2008 even with the default modeling scenario. Can the author explain why there is such a large difference in the month-to-month variation before and after 2008 in the default modeling scenario? Was it due to difference in NH3 emission inventory before and after 2008 or other reasons?

We explained the large MMV before 2008 in Section 3.2. The main idea is that when sulfate is high before 2008, the relatively low  $NH_3$  emission in August is insufficient relative to sulfate and leads to a much higher aerosol H<sup>+</sup> (than in June and July), which then leads to high IEPOX-SOA. After 2008 when sulfate becomes small, aerosol H<sup>+</sup> becomes less sensitive to the supply of  $NH_3$ , so even there are still monthly differences in  $NH_3$  emissions, the MMV of IEPOX-SOA is smaller.

In Section 2.2.2 At Line 301-303 we have the explanation: "Prior to 2008, IEPOX-SOA production is largely enhanced by abundant sulfate. Due to this high level of sulfate (about >4

 $\mu$ g/m<sup>3</sup>), the modeled aerosol acidity becomes particularly sensitive to variations in NH<sub>3</sub> emissions. ... The NH<sub>3</sub> emissions in August are about 25% lower than in June and July (Figure S4). As a result, in August before 2008, the aerosol NH<sub>4</sub><sup>+</sup>/SO<sub>4</sub><sup>2-</sup> ratio is smaller (Figure S4) and a<sub>H+</sub> is up to 3 times higher than June (Figure 4B), leading to high production of IEPOX-SOA in August. Both sulfate and aerosol acidity appear to be the dominant contributors to MMV of OA during this period. After 2008, IEPOX-SOA formation is substantially suppressed, due to small SO<sub>2</sub> emissions and low modeled aerosol acidity a<sub>H+</sub> with small monthly variability."

In addition, after updating the NH3 inventory with the new emission inventory from CrIS satellite observation, the author stated in line 384-385 that this scenario (CT\_newNH3) performed a bit worse than the fixed acidity scenario (CT\_H01). I wonder whether that was due to coating effects not correctly represented by using ethylene glycol rather than the results from authentic coatings. Besides the paper mentioned above, Li et al. 2020 and Zhang et al. also used simplified equations that can estimate the phase state of a few IEPOX-SOA species that might be helpful in performing future modeling.

Thank you for the suggestions. The CT\_newNH<sub>3</sub> simulation may be influenced by both the simplified coating scheme as well as uncertainties associated with NH<sub>3</sub> emissions. In the text we have:

"The CT\_newNH<sub>3</sub> scheme simulates an average IEPOX-SOA concentration of 0.34  $\mu$ g/m<sup>3</sup>, lower than the observation and the other models by a factor of >2, due to both the simplified coating effect and small aerosol  $a_{H^+}$  values ( $a_{H^+}$ <0.1mol/L, Figure 4B)." The Li et al. and Zhang et al. papers have been added to the model description Section 2.2.2 (see response to the second comment).

Minor comments: Why would there be a large increase of the default IEPOX-SOA during 2005-2008? The manuscript mentioned about higher correlation of IEPOX-SOA with acidity during this period of time. Was this abrupt increase of IEPOX-SOA caused by lower NH3 emissions between 2005-2008?

The high IEPOX-SOA in 2005-2007 is a result of both high sulfate, high aerosol acidity in August and high isoprene emissions.

In Section 3.2 at Line 314, we add: "The high IEPOX-SOA in 2000-2001 and 2005-2007 is a result of high sulfate aerosol, high aerosol acidity due to low NH<sub>3</sub> supply relative to high sulfate, and high isoprene emissions during these periods (Figure 3, Figure 4B)."

Line 171: There are multiple papers discussing about different Henry's law constants for IEPOX. The author did discuss in line 439 but probably should include other relevant papers, such as Woo et al., 2015, Budisulistiorini et al., 2016, Pye et al., 2017, and Zhang et al., 2018.

Thank you for the suggestions. We add at Line 439:

"the Henry's law constant for IEPOX uptake was tuned using measurements from the SOAS2013 campaign and was found to be  $1.7 \times 10^7$  M/atm, 10 times smaller than suggested by

Gaston et al. (2014) based on laboratory experiments and about half of the suggested value (3×10<sup>7</sup> M/atm) in some other studies (Budisulistiorini et al., 2017; Nguyen et al., 2014; Pye et al., 2017; Woo and McNeill, 2015; Zhang et al., 2018b)."

Line 477-478: One other potential mechanism I can think of is the non-linear feedback between sulfate and IEPOX-SOA production discussed in recent studies. For instance, Riva et al. 2019 and Zhang et al. both show that IEPOX-SOA fraction could and sulfate are nonlinear due to chemical reactions, acidity, and the coating effects of IEPOX-SOA are intertwined and nonlinear due to the formation of organosulfates.

Thank you for the suggestions. We mentioned the non-linearity at Line 408-410, and the Zhang et al. paper has been added as a reference now:

"Recent studies (Riva et al., 2019) suggested that the IEPOX-SOA production per unit mass of sulfate likely increases with decreasing sulfate due to changes in aerosol properties, such as acidity, morphology, phase state and viscosity, as well as formation of organosulfates, suggesting non-linearity between IEPOX-SOA and sulfate (Riva et al., 2019; Zhang et al., 2019)."

# **References:**

Budisulistiorini, S. H. et al. Seasonal characterization of submicron aerosol composition and organic aerosol sources in the southeastern United States: Atlanta, Georgia, and Look Rock, Tennessee. Atmos. Chem. Phys. 16, 5171-5189, doi:10.5194/acp-16-5171-2016 (2016).

Budisulistiorini, S. H. et al. Simulating Aqueous-Phase Isoprene-Epoxydiol (IEPOX) Secondary Organic Aerosol Production During the 2013 Southern Oxidant and Aerosol Study (SOAS). Environ. Sci. Technol. 51, 5026-5034, doi:10.1021/acs.est.6b05750 (2017).

Li, Y., Day, D. A., Stark, H., Jimenez, J. L. & Shiraiwa, M. Predictions of the glass transition temperature and viscosity of organic aerosols from volatility distributions. Atmos. Chem. Phys. 20, 8103-8122, doi:10.5194/acp-20-8103-2020 (2020).

Pye, H. O. T. et al. On the implications of aerosol liquid water and phase separation for organic aerosol mass. Atmos. Chem. Phys. 17, 343-369, doi:10.5194/acp17-343-2017 (2017).

Schmedding, R. et al.  $\alpha$ -Pinene-Derived organic coatings on acidic sulfate aerosol impacts secondary organic aerosol formation from isoprene in a box model. Atmos. Environ., 456-462, doi:https://doi.org/10.1016/j.atmosenv.2019.06.005 (2019).

Jo, D. S. et al. A simplified parameterization of isoprene-epoxydiol-derived secondary organic aerosol (IEPOX-SOA) for global chemistry and climate models: a case study with GEOS-Chem v11-02-rc. Geosci. Model Dev. 12, 2983-3000, doi:10.5194/gmd12-2983-2019 (2019).

Woo, J. L. & McNeill, V. F. simpleGAMMA v1.0 – a reduced model of secondary organic aerosol formation in the aqueous aerosol phase (aaSOA). Geosci. Model Dev. 8, 1821-1829, doi:10.5194/gmd-8-1821-2015 (2015).

Xu, L. et al. Effects of anthropogenic emissions on aerosol formation from isoprene and monoterpenes in the southeastern United States. Proc. Natl. Acad. Sci. USA 112, 37-42, doi:10.1073/pnas.1417609112 (2015).

Zhang, Y. et al. Effect of Aerosol-Phase State on Secondary Organic Aerosol Formation from the Reactive Uptake of Isoprene-Derived Epoxydiols (IEPOX). Environ. Sci. Technol. Lett. 5, 167-174, doi:10.1021/acs.estlett.8b00044 (2018).

Zhang, Y. et al. Joint Impacts of Acidity and Viscosity on the Formation of Secondary Organic Aerosol from Isoprene Epoxydiols (IEPOX) in Phase Separated Particles. ACS Earth and Space Chemistry 3, 2646-2658, doi:10.1021/acsearthspacechem.9b00209 (2019).

Zhang, Y. et al. The Cooling Rate- and Volatility-Dependent Glass-Forming Properties of Organic Aerosols Measured by Broadband Dielectric Spectroscopy. Environ. Sci. Technol. 53, 12366-12378, doi:10.1021/acs.est.9b03317 (2019).

# List of major changes in the manuscript

In this list, only major changes in the figures and main text are included. Other minor changes or rephrase of wording are included in the Responses to reviewers #1 and #2, and are shown in the attached manuscript with tracking changes. The page and line numbers refer to the revised version of the manuscript.

1. Another affiliation has been added to the author Eloise A. Marais.

2. Line 145, added the description of non-volatile cations and the thermodynamic model: "GEOS-Chem v12.1.1 considers sulfate, nitrate, and ammonium from all sectors, and fine-mode Na<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, Cl<sup>-</sup> from anthropogenic and sea salt sources, and employs the ISORROPIA II thermodynamic model (Fountoukis and Nenes, 2007; Pye et al., 2009; Song et al., 2018) to calculate aerosol water content and aerosol acidity (Pye et al., 2020)."

3. Line 159, added the description: "The default modeled OA shows a stronger decreasing trend from 2000 to 2013, and a large month-to-month variability in early 2000s, different from the observations (more details in Figure 1, 2A and Section 3.1). To address this model-observation discrepancy, we do four sets of  $2^{\circ} \times 2.5^{\circ}$  simulations: ... The sensitivity simulations are further explained in Section 2.2.2-2.2.3 and Section 3.3."

4. Line 187, added the description: "In the real atmosphere, inorganic aerosol is generally internally mixed with other organics. The presence of an organic coating may alter the aerosol properties and suppress the uptake of IEPOX onto acidified sulfate aerosol (Anttila et al., 2006; Gaston et al., 2014)."

5. Line 197, added the discussion of the uncertainties associated with coating parameterization: "In the real atmosphere when inorganic cores are coated with more viscous SOA (Zhang et al., 2018b), coating effect may be stronger because ethylene glycol is a low viscosity material. However, this simplified linear function does not consider the decreased viscosity and reduced coating effect at higher RH conditions (which is common in summertime southeast US) (Gaston et al., 2014; Zhang et al., 2018b), and prevents further IEPOX uptake when the mass fraction of OA ( $\chi_{org}$ ) is larger than 0.7, therefore this linear function may mimic a strong coating effect even though ethylene glycol is less viscous than real atmospheric SOA. The uncertainties need to be addressed in further studies with a more realistic coating parameterization (Li et al., 2020; Schmedding et al., 2019; Zhang et al., 2019b). We assume all OA is coated outside the inorganic aerosol core when calculating the IEPOX reactive uptake. The default GEOS-Chem with no organic coating effect, the increased particle radius  $R_p$  and surface area  $S_a$  of the mixed particle will partially offset (but does not outweigh) the impact of reduced reaction probability  $\gamma_{IEPOX.modified}$ ."

6. Line 220, added: "The default NEI11 and CrIS-derived NH<sub>3</sub> emissions averaged over 2000-2013 are compared in Figure S3." Line 225, added: "The CrIS-derived NH<sub>3</sub> emissions have been validated against surface observations of NH<sub>3</sub> concentration from the Ammonia Monitoring

Network (AMoN) and wet deposition measurements from the National Atmospheric Deposition Program (NADP). More details can be found in Cao et al. (2020)."

7. Line 281, added: "By sampling the model results at the locations of the IMPROVE and SEARCH sites, the modeled summertime OA has an average of 6.9  $\mu$ g/m<sup>3</sup> and a trend of 5.0%/year, similar to the model results averaged over the whole southeast US domain. For simplicity, we show only the domain-averaged model results in all figures and analysis."

8. Line 288, added: "The simulated IEPOX-SOA in 2013 compares well with previous field studies which suggested that IEPOX-SOA contributed to 18~40% in southeast US sites in summer 2013 (Budisulistiorini et al., 2016; Xu et al., 2015b)."

9. Line 333, added the temperature effect: "Temperature impacts the formation of IEPOX-SOA mainly through regulating isoprene emissions but does not influence partitioning as IEPOX-SOA is treated as non-volatile in GEOS-Chem. Therefore, temperature is not examined as another driver in addition to isoprene emissions."

10. Line 343, added the description of Figure 3: "Figure 3 shows the standardized monthly surface IEPOX-SOA concentration, sulfate concentration, aerosol  $H^+$  activity and isoprene emission from the default model. For each variable, the monthly gridded data has been first averaged over the southeast US. Then, we calculate the one standard deviation of all monthly data (June, July and August data from 2000 to 2013). Finally, the domain-averaged monthly data has been divided by its standard deviation, so the variables are standardized to be unitless and their variability can be compared directly."

11. Line 351, modified to: "The default NH<sub>3</sub> emissions from NEI11v1 suggest no significant long-term trend from 2000 to 2013. In general, ammonium aerosol is strongly correlated with sulfate and has a similar declining trend as sulfate (Silvern et al., 2017). However, the NH<sub>3</sub> emissions in August are about 25% lower than in June and July (Figure S4). As a result, in August before 2008, the aerosol NH<sub>4</sub><sup>+</sup>/SO<sub>4</sub><sup>2-</sup> ratio is smaller (Figure S4) and  $a_{H^+}$  is up to 3 times higher than June (Figure 4B), leading to high production of IEPOX-SOA in August."

12. Line 371, modified to: "Isoprene emission contributes to the overall interannual variability, for example leading to the relatively low IEPOX-SOA in 2003-2004 and the peaks in 2000, 2006 and 2011 (Figure 3). The high IEPOX-SOA in 2000-2001 and 2005-2007 is a result of high sulfate aerosol, high aerosol acidity due to low NH<sub>3</sub> supply relative to high sulfate, and high isoprene emissions during these periods (Figure 3, Figure 4B)."

13. Line 460, added: "Recent studies (Riva et al., 2019) suggested that the IEPOX-SOA production per unit mass of sulfate likely increases with decreasing sulfate due to changes in aerosol properties, such as acidity, morphology, phase state and viscosity, **as well as formation of organosulfates, suggesting non-linearity between IEPOX-SOA and sulfate (Riva et al., 2019; Zhang et al., 2019a).**"

14. Line 490, added: "In the default scheme, the Henry's law constant for IEPOX uptake was tuned using measurements from the SOAS2013 campaign and was found to be  $1.7 \times 10^7$  M/atm,

10 times smaller than suggested by Gaston et al. (2014) based on laboratory experiments and about half of the suggested value (3×10<sup>7</sup> M/atm) in some other studies (Budisulistiorini et al., 2017; Nguyen et al., 2014; Pye et al., 2017; Woo and McNeill, 2015; Zhang et al., 2018b)."

15. Line 513, added discussion on the limitation of using a coarse horizontal resolution: "Simulations in this study are conducted at a horizontal resolution of  $2^{\circ} \times 2.5^{\circ}$ , which is comparable to most global climate models. However, as shown by Yu et al. (2016), from coarse to fine horizontal resolution, there will be a shift from low-NO<sub>x</sub> to high-NO<sub>x</sub> pathway for isoprene oxidation. Therefore, using a fine resolution may reduce the production of IEPOX and IEPOX-SOA, which needs further investigation."

16. Line 524, added: "Such discrepancies may suggest a more important role of SOA pathways that are less dependent on inorganic aerosol, such as terpene-SOA formed by reversible gasaerosol partitioning. Terpene-SOA is included in GEOS-Chem (yellow color in Figure 1), and contributes to 8~24% of total OA, which might be underestimated compared to recent field studies."

17. Figure 2 has been revised to add the trend lines of organic aerosol and sulfate.

18. Figure S1 has been revised to add the location of CSN sites.

19. Added Figure S3 to compare the two NH<sub>3</sub> emission inventories.

20. Figure S4 (previous Figure S3) has been revised to add the trend of aerosol ammonium-to-sulfate ratio.

# Long-term observational constraints of organic aerosol dependence on inorganic species in the southeast US

Yiqi Zheng<sup>1,2,\*</sup>, Joel A. Thornton<sup>3</sup>, Nga Lee Ng<sup>4,5,6</sup>, Hansen Cao<sup>7</sup>, Daven K. Henze<sup>7</sup>, Erin E. McDuffie<sup>8,9</sup>, Weiwei Hu<sup>10,11</sup>, Jose L. Jimenez<sup>11</sup>, Eloise A. Marais<sup>12,13</sup>, Eric Edgerton<sup>14</sup>, Jingqiu Mao<sup>1,2,\*</sup>

<sup>1</sup>Geophysical Institute, University of Alaska Fairbanks, Fairbanks, AK, USA

<sup>2</sup>Department of Chemistry and Biochemistry, University of Alaska Fairbanks, Fairbanks, AK, USA

<sup>3</sup>Department of Atmospheric Sciences, University of Washington, Seattle, WA, USA

<sup>4</sup>School of Chemical and Biomolecular Engineering, Georgia Institute of Technology, Atlanta, GA, USA

<sup>5</sup>School of Earth and Atmospheric Sciences, Georgia Institute of Technology, Atlanta, GA, USA <sup>6</sup>School of Civil and Environmental Engineering, Georgia Institute of Technology, Atlanta, GA, USA

<sup>7</sup>Department of Mechanical Engineering, University of Colorado Boulder, Boulder, CO, USA <sup>8</sup>Department of Physics and Atmospheric Science, Dalhousie University, Halifax, Nova Scotia, Canada

<sup>9</sup>Department of Energy, Environmental & Chemical Engineering, Washington University in St. Louis, St Louis, MO, USA

<sup>10</sup>State Key Laboratory of Organic Geochemistry, Guangzhou Institute of Geochemistry, Chinese Academy of Science (CAS), Guangzhou, China

<sup>11</sup>Department of Chemistry and CIRES, University of Colorado Boulder, Boulder, CO, USA
 <sup>12</sup>School of Physics and Astronomy, University of Leicester, Leicester, LE1 7RH, UK
 <sup>13</sup>Now at Department of Geography, University College London, London, UK
 <sup>14</sup>Atmospheric Research & Analysis, Inc., Cary, NC, USA

\*Correspondence to: Yiqi Zheng (yzheng4@alaska.edu) and Jingqiu Mao (jmao2@alaska.edu)

1

Keywords: Secondary organic aerosol; IEPOX; isoprene; coating; aerosol acidity.

### 1 Abstract

2 Organic aerosol (OA), with a large biogenic fraction in summertime southeast US, adversely 3 impacts on air quality and human health. Stringent air quality controls have recently reduced anthropogenic pollutants including sulfate, whose impact on OA remains unclear. Three filter 4 5 measurement networks provide long-term constraints on the sensitivity of OA to changes in 6 inorganic species, including sulfate and ammonia. The 2000-2013 summertime OA decreases by 7 1.7~1.9%/year with little month-to-month variability, while sulfate declines rapidly with 8 significant monthly difference in early 2000s. In contrast, modeled OA from a chemical-9 transport model (GEOS-Chem) decreases by 4.9%/year with much larger monthly variability, 10 largely due to the predominant role of acid-catalyzed reactive uptake of epoxydiols (IEPOX) onto sulfate. The overestimated modeled OA dependence on sulfate can be improved by 11 12 implementing a coating effect and assuming constant aerosol acidity, suggesting the needs to 13 revisit IEPOX reactive uptake in current models. Our work highlights the importance of 14 secondary OA formation pathways that are weakly dependent on inorganic aerosol in a region 15 that is heavily influenced by both biogenic and anthropogenic emissions. 16

17

### 18 1. Introduction

- 19 Organic aerosol (OA) accounts for a large fraction of ambient fine particulate matter mass, 20 which strongly affects air quality, regional climate, and human welfare (Jimenez et al., 2009). Since the implementation of the Clean Air Act Amendments of 1990, there has been a significant 21 22 decline in ambient aerosol in the United States, mostly due to reductions in inorganic aerosol 23 mass following changes in emissions of sulfur dioxide (SO<sub>2</sub>), nitrogen oxides (NO<sub>x</sub> =NO+ NO<sub>2</sub>), 24 as well as reductions in black carbon (EPA, 2011), leaving OA as the major component of fine particulate matter (50~70%) over the southeast US, especially in summer (Attwood et al., 2014; 25 Kim et al., 2015). OA can be directly emitted by combustion processes (primary organic aerosol, 26 27 POA) or secondarily formed (SOA) from the atmospheric oxidation of biogenic volatile organic 28 compounds (BVOCs), mainly isoprene and monoterpenes, and also precursors from 29 anthropogenic sources and biomass burning (Hayes et al., 2015; Hodshire et al., 2019). OA has 30 also been declining across much of the US over the past few decades, primarily due to decreased 31 anthropogenic emissions from vehicle and residential fuel-burning, except for the southeast US 32 (Ridley et al., 2018). The southeast US is one of the largest BVOC emission hotspots in the 33 world (Guenther et al., 2006), and at the same time is heavily populated with large anthropogenic 34 emissions of pollutants. Biogenic SOA (formed from atmospheric oxidation of BVOCs) may 35 account for 60-100% of OA in summertime southeast US (Kim et al., 2015; Xu et al., 2015b). To what extent biogenic SOA could be mediated through emission control strategies remains an 36 37 open question (Carlton et al., 2010; Mao et al., 2018). 38
- The oxidation of BVOCs produces hundreds of intermediate products. Some products have low
  volatility that can partition onto the condensed phase, while some gas-phase products can react in

41	the aqueous phase to form SOA. SOA formed from uptake of isoprene epoxydiols (IEPOX-
42	SOA) (Paulot et al., 2009) appears to be the major confirmed aqueous SOA product globally,
43	being important in all high isoprene and lower NO regions (Hu et al., 2015), along with glyoxal
44	formed from isoprene and aromatics (Fu et al., 2008). Formation of SOA in clouds was
45	investigated in the southeast US and found to be not statistically significant (Wagner et al.,
46	2015). These pathways have been implemented into 3-dimensional global atmospheric chemistry
47	and climate models using two different approaches. First, to simulate the partitioning of organic
48	vapors, the BVOC oxidation products can be grouped based on their volatility (Volatility Basis
49	Set, VBS), and the product yields and vapor pressures are parameterized for each surrogate
50	precursor (Donahue et al., 2006; Pankow, 1994). Such empirical VBS schemes are usually
51	derived using dry laboratory chamber experiments (with relative humidity $RH < 10\%$ ) and do not
52	explicitly depend on aerosol water, RH, or inorganic aerosol mass or composition. Therefore,
53	here we refer to the SOA formed through partitioning calculated by VBS as dry SOA. Second, a
54	more explicit representation of aqueous SOA formation from isoprene products has been used
55	recently, which incorporates dependence on inorganic aerosol volume and aerosol acidity
56	(Budisulistiorini et al., 2017; Ervens et al., 2011; Fu et al., 2008; Marais et al., 2016; Pye et al.,
57	2013). The relative contribution of dry versus aqueous SOA to total OA mass in the atmosphere
58	is uncertain and has limited observational constraints.
59	

Long-term field measurements show a decreasing trend of OA in the southeast US (Attwood et
al., 2014; Hidy et al., 2014; Kim et al., 2015), which is likely linked to reductions in
anthropogenic POA and SOA (Blanchard et al., 2016; Ridley et al., 2018), sulfate (Blanchard et
al., 2016; Malm et al., 2017; Marais et al., 2017; Xu et al., 2015b, 2016) and NO<sub>x</sub> (Carlton et al.,

64	2010; Pye et al., 2010, 2019; Xu et al., 2015b). The influence of sulfate on OA is thought to be
65	mainly due to its influence on the uptake of isoprene gas-phase oxidation products, which are
66	often small molecules that cannot directly condense due to high vapor pressure, but may undergo
67	aqueous-phase reactive uptake onto wet sulfate particles to form aqueous SOA, as suggested by
68	extensive laboratory and field studies (Budisulistiorini et al., 2015; Hu et al., 2015; Li et al.,
69	2016; Liggio et al., 2005; McNeill et al., 2012; Riedel et al., 2016; Shrivastava et al., 2017;
70	Surratt et al., 2010; Tan et al., 2012; Xu et al., 2016, 2015b). NOx plays a complex role in
71	regulating oxidation capacity, different oxidation pathways and aerosol water content through
72	aerosol nitrate (Kiendler-Scharr et al., 2016; Kroll et al., 2005, 2006; Li et al., 2018; Ng et al.,
73	2017; Presto et al., 2005; Shrivastava et al., 2019; Zheng et al., 2015; Ziemann and Atkinson,
74	2012). Prior 3-D modeling studies with different SOA mechanisms provide different
75	explanations for the long-term OA trend observed in the southeast US. For example, the dry
76	SOA calculated by VBS framework with NOx-dependent yields implies a small decrease in OA
77	following the reductions of $NO_x$ (Pye et al., 2013; Zheng et al., 2015), but has little dependence
78	on changes in inorganic aerosol mass such as sulfate. On the other hand, models using aqueous
79	SOA formation from isoprene attributed the decreasing OA from 1991 to 2013 to reductions in
80	sulfate (Marais et al., 2017) but showed greater interannual variability than was observed. The
81	driving mechanism for the OA trend in the southeast US remains to be elucidated.
82	

- 83 Here we use observations from three surface filter-based networks (IMPROVE, SEARCH,
- 84 CSN), combined with a 3-dimensional chemical transport model GEOS-Chem v12.1.0, to
- 85 examine the long-term trend and more importantly, the month-to-month variability of OA in the

86 southeast US during 2000-2013. The results provide new observational constraints on the drivers

87 of OA variability and the SOA formation mechanisms in the southeast US.

88

#### 89 2. Methods

#### 90 2.1 Observational datasets.

91 We use surface filter-based measurement of fine particulate matter mass and composition

92 (including organic carbon, OC) in 2000-2013 from three networks: the Interagency Monitoring

93 of Protected Visual Environments (IMPROVE) (Solomon et al., 2014), the SouthEastern Aerosol

94 Research and Characterization (SEARCH) (Edgerton et al., 2005), and the Environmental

95 Protection Agency's PM2.5 National Chemical Speciation Network (CSN) (Solomon et al.,

96 2014). We select 21 IMPROVE sites, 3 SEARCH rural sites and 36 CSN sites within the

97 southeast US region [29°~37°N, 74°~96°W] (Figure S1). The SEARCH sites are organized in

98 rural/urban pairs (Edgerton et al., 2005) and only the data from the rural sites are used here to

99 represent background conditions. IMPROVE sites are mostly rural (Solomon et al., 2014). The

100 OC measurement in the CSN network in 2004-2009 gradually shifted to a different protocol and

101 analytical technique than the early 2000s, which led to the discontinuity in long-term trend

102 (Figure S2), therefore we only use CSN data to examine the monthly variability of OA, and

103 focus on IMPROVE and SEARCH for all analysis. The 3-day OC measurement from IMPROVE

104 and daily OC from SEARCH and CSN are averaged to monthly values. A factor of 2.1 is used to

105 convert measured organic carbon (OC) to organic aerosol mass, as suggested by the southeast

106 US field measurements (Pye et al., 2017; Schroder et al., 2018).

107

108	We use OA measurements by Aerosol Mass Spectrometer (AMS) from the Southern Oxidant	
109	and Aerosol Studies campaign (SOAS) at the Centerville, AL Site in 06/01/2013-07/15/2013	
110	(SOAS2013). The OA measurements and derived IEPOX-SOA factor calculated by Positive	
111	Matrix Factorization (PMF) analysis (Hu et al., 2015; Xu et al., 2015a, 2018) are from two	
112	independent groups: one group from Georgia Institute of Technology led by Prof. Nag Lee Ng,	
113	the other from University of Colorado Boulder led by Prof. Jose L. Jimenez, denoted as Obs_GT	
114	and Obs_CU, respectively.	
115		
116	2.2 Modeling framework	
117	2.2.1 GEOS-Chem	
118	In this study we use the 3-dimensional global chemical transport model GEOS-Chem version	
119	12.1.1 (DOI: 10.5281/zenodo.2249246, https://github.com/geoschem/) with detailed O3-NOx-	
120	HOx-CO-VOC-aerosol tropospheric chemistry (Bey et al., 2001; Mao et al., 2013). Isoprene	
121	chemistry is described in (Fisher et al., 2016; Travis et al., 2016). GEOS-Chem is driven by	
122	offline meteorology 1999-2013 from the NASA Modern-Era Retrospective analysis for Research	
123	and Applications, version 2 (MERRA-2 https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/). The	
124	global anthropogenic (including agricultural) emissions are from the Community Emissions Data	
125	System (CEDS) inventory, with the US region replaced by the EPA's National Emission	
126	Inventory for 2011 (NEI11v1). The monthly mean anthropogenic emissions of CO, SO <sub>2</sub> , NO <sub>x</sub> ,	
127	NH <sub>3</sub> , VOCs, OC and black carbon are scaled to the year 2011 using the ratio of EPA's national	
128	annual emission totals from 2000 to 2013 (Travis et al., 2016). Biomass burning emissions are	
129	from Global Fire Emissions Database version 4 (GFED4) (Randerson et al., 2015). Biogenic	

130 emissions of isoprene and terpenes are online calculated by the Model of Emissions of Gases and

**Deleted:** mapped over 0.1°×0.1° and

132Aerosols from Nature (MEGAN2.1) (Guenther et al., 2012) that is also driven by MERRA-2

133 meteorology.

134

135	For organic aerosol, we employ the complexSOA scheme for SOA modeling for all simulations
136	in this study (Marais et al., 2016; Pai et al., 2020; Pye et al., 2010). POA are regarded as
137	nonvolatile. This SOA modeling includes a 4-product Volatility-Basis-Set (VBS) for SOA
138	formation from reversible condensation of oxidation products of biogenic terpenes (including
139	monoterpenes and sesquiterpenes), and anthropogenic VOCs, referred to as terpene-SOA and
140	anthropogenic SOA, respectively. The SOA calculated through VBS parameterization is fitted
141	based on dry chamber (RH<10%) results independent of inorganic aerosol, aerosol water and RH
142	(Pye et al., 2010). The complexSOA scheme also includes aqueous SOA formed from reactive
143	uptake of isoprene oxidation products, including IEPOX, glyoxal, C4 epoxides, methylglyoxal,
144	non-IEPOX product of the ISOPOOH oxidation and hydroxynitrates from NO3-initiated
145	oxidation (Marais et al., 2016). GEOS-Chem v12.1.1 considers sulfate, nitrate, and ammonium
146	from all sectors, and fine-mode Na <sup>+</sup> , Ca <sup>2+</sup> , Mg <sup>2+</sup> , Cl <sup>-</sup> from anthropogenic and sea salt sources,
147	and employs the ISORROPIA II thermodynamic model (Fountoukis and Nenes, 2007; Pye et al.,
148	2009; Song et al., 2018) to calculate aerosol water content and aerosol acidity (Pye et al., 2020).
149	
150	We run the <u>default</u> GEOS-Chem model at $4^{\circ} \times 5^{\circ}$ latitude by longitude continuously from
151	10/01/1999 to 12/31/2013. For each year, the restart file at 05/01 from the continuous $4^{\circ} \times 5^{\circ}$
152	simulation has been regridded to $2^{\circ} \times 2.5^{\circ}$ and is used to initiate $2^{\circ} \times 2.5^{\circ}$ simulations from $05/01$
153	to 08/31 each year. The May results are discarded as spin-up and the results of June, July and
154	August are used for analysis.

**Deleted:** The  $2^{\circ} \times 2.5^{\circ}$  simulations are adequate when modeling continental boundary layer chemistry (Yu et al., 2016).

158	
159	The default modeled OA shows a stronger decreasing trend from 2000 to 2013, and a large
160	month-to-month variability in early 2000s, different from the observations (more details in
161	Figure 1, 2A and Section 3.1). To address this model-observation discrepancy, we do four sets of
162	2°×2.5° simulations: Default (using default complexSOA scheme); CT (with coating effect for
163	IEPOX reactive uptake); CT_newNH3 (with coating effect and US NH3 emissions replaced by
164	satellite-derived NH <sub>3</sub> inventory); CT_H01 (with coating effect and fixing aerosol <u>acidity <math>a_{H+}</math> at</u>
165	0.1 mol/L when calculating IEPOX reactive uptake). The sensitivity simulations are further
166	explained in Section 2.2.2-2.2.3 and Section 3.3.

167

#### 168 **2.2.2 Coating**

The default IEPOX-SOA mechanism in GEOS-Chem uses aerosol-phase reaction rates from
laboratory chamber studies with pure acidic inorganic particles (Gaston et al., 2014; Riedel et al.,
2015), and a representative effective Henry's law constant obtained by matching the model to the
observations from the SOAS2013 campaign\_(Marais et al., 2016), to estimate the reactive uptake
coefficient *piEPOX*. In the default scheme, *piEPOX* is calculated as follows:

174 
$$\frac{1}{\gamma_{IEPOX}} = \frac{R_p \omega}{4D_g} + \frac{1}{\alpha} + \frac{1}{\Gamma_{aq}}$$

175 
$$\Gamma_{aq} = \frac{4VRTH_{aq}k_{aq}}{S_a\omega}$$

176 Where  $R_p$  is the particle radius of the inorganic sulfate-nitrate-ammonium particle (cm),  $\omega$  is the 177 mean molecular speed (cm/s),  $D_g$  is the gas-phase diffusion coefficient (0.1 cm<sup>2</sup>/s),  $\alpha$  is the mass 178 accommodation coefficient ( $\alpha$ =0.1),  $S_a$  is the total (wet) particle surface area (cm<sup>2</sup>/cm<sup>3</sup>), V is the 179 total (wet) particle volume (cm<sup>3</sup>/cm<sup>3</sup>), R is the ideal gas constant (L atm/mol/K), T is temperature 180 (K),  $H_{aq}$  is the Henry's law coefficient (1.7×10<sup>7</sup> M/atm), and  $k_{aq}$  is the first-order reaction rate 181 constant  $(s^{-1})$ :  $k_{aq} = k_{H^+}[H^+] + k_{nuc}[nuc]a_{H^+} + k_{ga}[ga]$ 182 where  $k_{H^+}$  (=0.036 M<sup>-1</sup>s<sup>-1</sup>),  $k_{nuc}$  (=2×10<sup>-4</sup> M<sup>-1</sup>s<sup>-1</sup>) and  $k_{ga}$  (=7.3×10<sup>-4</sup> M<sup>-1</sup>s<sup>-1</sup>) are the reaction 183 184 rates due to acid-catalyzed ring-opening, presence of nucleophiles (including nitrate and sulfate) 185 and presence of bisulfate acids, respectively (Gaston et al., 2014; Marais et al., 2016). 186 187 In the real atmosphere, inorganic aerosol is generally internally mixed with other organics. The 188 presence of an organic coating may alter the aerosol properties and suppress the uptake of 189 IEPOX onto acidified sulfate aerosol (Anttila et al., 2006; Gaston et al., 2014). We implement a 190 linear coating effect for the IEPOX-SOA formation. The coating effect is fitted using laboratory-191 derived values of *y*<sub>IEPOX</sub> on particles containing both ammonium bisulfate and ethylene glycol 192 under RH=50% conditions (Gaston et al., 2014). In the coating scheme,  $\gamma'_{IEPOX}$  is calculated as 193 above with  $R_p$ , V and  $S_a$  updated considering OA coated outside the inorganic core. Then, the 194 fitted function is applied to modify  $\gamma'_{IEPOX}$ : 195  $\gamma_{IEPOX\_modified} = \gamma'_{IEPOX} \times (1 - 1.3 \times \chi_{org})$ 196 where  $\chi_{org}$  is the mass fraction of OA in the mixed particle including both the inorganic aerosol 197 and OA. When  $\chi_{org} > 0.7$ , the IEPOX uptake will be terminated, i.e.  $\gamma_{IEPOX\_modified} = 0$ . In the 198 real atmosphere when inorganic cores are coated with more viscous SOA (Zhang et al., 2018b), 199 coating effect may be stronger because ethylene glycol is a low viscosity material. However, this simplified linear function does not consider the decreased viscosity and reduced coating effect at 200 201 higher RH conditions (which is common in summertime southeast US) (Gaston et al., 2014; 202 Zhang et al., 2018b), and prevents further IEPOX uptake when the mass fraction of OA ( $\chi_{org}$ ) is

203	larger than 0.7, therefore this linear function may mimic a strong coating effect even though
204	ethylene glycol is less viscous than real atmospheric SOA. The uncertainties need to be
205	addressed in further studies with a more realistic coating parameterization (Li et al., 2020;
206	Schmedding et al., 2019; Zhang et al., 2019b). We assume all OA is coated outside the inorganic
207	aerosol core when calculating the IEPOX reactive uptake. The default GEOS-Chem with no
208	organic coating calculates surface area of inorganic aerosol (Jo et al., 2019). By adding the
209	<u>coating effect</u> , the increased particle radius $R_p$ and surface area $S_a$ of the mixed particle will
210	partially offset (but does not outweigh) the impact of reduced reaction probability
211	YIEPOX_modified:
212	
213	2.2.3 Satellite-derived NH <sub>3</sub> emissions
214	We use the Cross-track Infrared Sounder (CrIS) satellite-derived NH3 emissions_(Cao et al.,
215	2020) in a sensitivity test in this study. The top-down monthly NH <sub>3</sub> emissions over the
216	contiguous US at $0.25^{\circ} \times 0.3125^{\circ}$ latitude by longitude are derived from CrIS v1.5
217	measurements of $NH_3$ profiles (Shephard and Cady-Pereira, 2015) for the year 2014 through a
218	4D-Var approach using GEOS-Chem and its adjoint model (Henze et al., 2007). The CrIS-
219	derived emissions are then regridded to $0.1^{\circ} \times 0.1^{\circ}$ to replace the default NEI11 emissions for the
220	year 2011 and applied the same annual scaling factors for 2000-2013. The default NEI11 and
221	CrIS-derived NH <sub>3</sub> emissions averaged over 2000-2013 are compared in Figure S3. There is no
222	significant trend <u>of <math>NH_3</math> emissions</u> from 2000 to 2013 (Figure S <u>4</u> ), consistent with other studies
223	suggesting nearly constant NH <sub>3</sub> emissions from 2001 to 2014 (Butler et al., 2016). The CrIS-
224	derived emissions used the HTAPv2 emissions inventory as the prior emissions, which is based
225	on the 2008 NEI emissions over the US (Janssens-Maenhout et al., 2015). The CrIS-derived NH <sub>3</sub>
1	

226	emissions have been validated against surface observations of $NH_3$ concentration from the	
227	Ammonia Monitoring Network (AMoN) and wet deposition measurements from the National	
228	Atmospheric Deposition Program (NADP). More details can be found in Cao et al. (2020). Using	<b>Deleted:</b> The top-down annual mean emiss
229	the top-down emissions in GEOS-Chem increases the correlation coefficient $(r)$ between	emissions underestimate agricultural emiss springtime fertilizer and livestock sources of
230	modeled monthly mean $NH_3$ concentration and surface observations from 0.74 to 0.93 and	emissions were found in the Central Valley Minnesota, northern Iowa and southeast No
231	reduces the normalized mean bias of domain-averaged annual mean simulated NH3 by a factor of	during warm months.
232	1.9. The seasonal cycle of simulated wet $NH_4^+$ deposition is also improved ( <i>r</i> increased from	
233	0.70 to 0.86), but the normalized mean bias of domain-averaged annual simulated wet $\mathrm{NH_4^+}$	
234	increases from 0.34 to 0.96 due to overly strong wet scavenging in the model. The latter issue	
235	was ultimately resolved in Cao et al. (2020) and the final top-down emissions reported therein	
236	differ from those reported here; nevertheless, the emissions estimates used here provide a	
237	valuable basis for conducting a sensitivity experiment,	Formatted: Font color: Auto, Not Raised b
237 238	valuable basis for conducting a sensitivity experiment,	Formatted: Font color: Auto, Not Raised I
237 238 239	valuable basis for conducting a sensitivity experiment, 2.3 Multivariate linear regression analysis	Formatted: Font color: Auto, Not Raised l
237 238 239 240	valuable basis for conducting a sensitivity experiment, <b>2.3 Multivariate linear regression analysis</b> In this study we did a multivariate regression analysis of modeled monthly IEPOX-SOA (μg/m <sup>3</sup> )	Formatted: Font color: Auto, Not Raised l
237 238 239 240 241	valuable basis for conducting a sensitivity experiment, <b>2.3 Multivariate linear regression analysis</b> In this study we did a multivariate regression analysis of modeled monthly IEPOX-SOA ( $\mu$ g/m <sup>3</sup> ) against modeled sulfate aerosol ( $\mu$ g/m <sup>3</sup> ), aerosol acidity $a_{H^+}$ (mol/L) and isoprene emission	Formatted: Font color: Auto, Not Raised I
237 238 239 240 241 242	valuable basis for conducting a sensitivity experiment, <b>2.3 Multivariate linear regression analysis</b> In this study we did a multivariate regression analysis of modeled monthly IEPOX-SOA ( $\mu$ g/m <sup>3</sup> ) against modeled sulfate aerosol ( $\mu$ g/m <sup>3</sup> ), aerosol acidity $a_{H^+}$ (mol/L) and isoprene emission ( <i>ISOP</i> <sub>emis</sub> mg/m <sup>2</sup> /hr):	Formatted: Font color: Auto, Not Raised I
<ul> <li>237</li> <li>238</li> <li>239</li> <li>240</li> <li>241</li> <li>242</li> <li>243</li> </ul>	valuable basis for conducting a sensitivity experiment, <b>2.3 Multivariate linear regression analysis</b> In this study we did a multivariate regression analysis of modeled monthly IEPOX-SOA ( $\mu$ g/m <sup>3</sup> ) against modeled sulfate aerosol ( $\mu$ g/m <sup>3</sup> ), aerosol acidity $a_{H^+}$ (mol/L) and isoprene emission ( <i>ISOP<sub>emis</sub></i> mg/m <sup>2</sup> /hr): <i>IEPOX-SOA</i> = $\beta_1 \times sulfate + \beta_2 \times a_{H^+} + \beta_3 \times ISOP_{emis} + b$	Formatted: Font color: Auto, Not Raised I
237 238 239 240 241 242 243 244	valuable basis for conducting a sensitivity experiment, <b>2.3 Multivariate linear regression analysis</b> In this study we did a multivariate regression analysis of modeled monthly IEPOX-SOA ( $\mu$ g/m <sup>3</sup> ) against modeled sulfate aerosol ( $\mu$ g/m <sup>3</sup> ), aerosol acidity $a_{H^+}$ (mol/L) and isoprene emission ( $ISOP_{emis}$ mg/m <sup>2</sup> /hr): $IEPOX-SOA = \beta_1 \times sulfate + \beta_2 \times a_{H^+} + \beta_3 \times ISOP_{emis} + b$ Mean values have been subtracted from all variables, which are then divided by standard	Formatted: Font color: Auto, Not Raised I
237 238 239 240 241 242 243 244 245	valuable basis for conducting a sensitivity experiment, <b>2.3 Multivariate linear regression analysis</b> In this study we did a multivariate regression analysis of modeled monthly IEPOX-SOA ( $\mu$ g/m <sup>3</sup> ) against modeled sulfate aerosol ( $\mu$ g/m <sup>3</sup> ), aerosol acidity $a_{H^+}$ (mol/L) and isoprene emission ( $ISOP_{emis}$ mg/m <sup>2</sup> /hr): $IEPOX-SOA = \beta_1 \times sulfate + \beta_2 \times a_{H^+} + \beta_3 \times ISOP_{emis} + b$ Mean values have been subtracted from all variables, which are then divided by standard deviations. $\beta_1$ , $\beta_2$ and $\beta_3$ are standardized partial regression coefficients associated with sulfate	Formatted: Font color: Auto, Not Raised I
<ul> <li>237</li> <li>238</li> <li>239</li> <li>240</li> <li>241</li> <li>242</li> <li>243</li> <li>244</li> <li>245</li> <li>246</li> </ul>	valuable basis for conducting a sensitivity experiment, <b>2.3 Multivariate linear regression analysis</b> In this study we did a multivariate regression analysis of modeled monthly IEPOX-SOA ( $\mu$ g/m <sup>3</sup> ) against modeled sulfate aerosol ( $\mu$ g/m <sup>3</sup> ), aerosol acidity $a_{H^+}$ (mol/L) and isoprene emission ( $ISOP_{emis}$ mg/m <sup>2</sup> /hr): $IEPOX-SOA = \beta_1 \times sulfate + \beta_2 \times a_{H^+} + \beta_3 \times ISOP_{emis} + b$ Mean values have been subtracted from all variables, which are then divided by standard deviations. $\beta_1$ , $\beta_2$ and $\beta_3$ are standardized partial regression coefficients associated with sulfate aerosol, $a_{H^+}$ and isoprene emission, and can be directly compared to evaluate the relative	Formatted: Font color: Auto, Not Raised I

importance of the three variables. We apply the regression analysis using monthly data within 247

sions are ~52% ause the prior sions, in particular over the Central the prior y, southern lorth Carolina

by / Lowered by

256 different time frames (2000-2013, 2000-2004, 2005-2008 and 2009-2013 as in Table S1) to

257 determine the evolving importance of variables.

258

259

## 260 **3. Results**

#### 261 3.1 Long-term trend and month-to-month variability (MMV) of OA

262 In the southeast US, observations from the IMPROVE and SEARCH network both show a

263 reduction in summertime surface OA concentration from 2000 to 2013 (Figure 1). Observational

264 results are averaged using 21 IMPROVE sites and 3 SEARCH sites within the southeast US. OA

265 concentration averaged over June-July-August (JJA) 2000-2013 is 4.2 µg/m<sup>3</sup> from the

266 IMPROVE sites, and 5.7  $\mu$ g/m<sup>3</sup> from SEARCH sites. A similar ~30% summertime low bias on

267 the IMPROVE sites was documented by Kim et al. (2015) compared to the SEARCH sites,

268 which is thought to be due to evaporation of OA from the filters after collection, as the

269 IMPROVE filters stay several days on site after sampling and are shipped without refrigeration,

270 while the SEARCH filters are analyzed in-situ. Despite different magnitudes, OA from the two

271 networks demonstrate similar trends and interannual variability. The 2000-2013 trend of JJA OA

272 mass is -1.7%/year for IMPROVE and -1.9%/year for SEARCH. Compared to the slow decrease

- 273 in OA, a faster declining trend is found for sulfate from IMPROVE (-6.9%/year) and SEARCH
- (-6.7%/year) for the same period (Figure 2).
- 275

276 Compared to the observations, the default GEOS-Chem model predicts a steeper decreasing

trend of OA mass during 2000-2013 (Figure 1). Modeling results are averaged over the domain

278 [29°~37°N, 74°~96°W] excluding ocean grid cells (Figure S1). The 2000-2013 JJA-averaged

279	OA from the default model is 6.7 $\mu$ g/m <sup>3</sup> , higher than OA from IMPROVE and SEARCH.
280	Modeled total OA mass decreases at a rate of 4.9%/year, about 1.9 (1.6) times faster than
281	IMPROVE (SEARCH) OA (student's t-test p<0.001). By sampling the model results at the
282	locations of the IMPROVE and SEARCH sites, the modeled summertime OA has an average of
283	$6.9 \mu g/m^3$ and a trend of 5.0%/year, similar to the model results averaged over the whole
284	southeast US domain. For simplicity, we show only the domain-averaged model results in all
285	figures and analysis. The strong reduction in total OA mass is dominated by aqueous SOA,
286	especially through reactive uptake of IEPOX, with no decreasing trend in other components
287	(Figure 1). The contribution of IEPOX-SOA to total OA mass decreases from 61% in the early
288	2000s to 28% in 2013. The simulated IEPOX-SOA in 2013 compares well with previous field
289	studies which suggested that IEPOX-SOA contributed to 18~40% in southeast US sites in
290	summer 2013 (Budisulistiorini et al., 2016; Xu et al., 2015b).
291	
292	A main constraint comes from the MMV of OA in the southeast US. IMPROVE and SEARCH
292 293	A main constraint comes from the MMV of OA in the southeast US. IMPROVE and SEARCH OA observations show little variability among June, July and August, despite large MMV of
292 293 294	A main constraint comes from the MMV of OA in the southeast US. IMPROVE and SEARCH OA observations show little variability among June, July and August, despite large MMV of sulfate in early 2000s (Figure 2A). We find similar behavior from another observation network,
292 293 294 295	A main constraint comes from the MMV of OA in the southeast US. IMPROVE and SEARCH OA observations show little variability among June, July and August, despite large MMV of sulfate in early 2000s (Figure 2A). We find similar behavior from another observation network, CSN. The discontinuity in OA trend in the CSN network is due to different protocols applied
292 293 294 295 296	A main constraint comes from the MMV of OA in the southeast US. IMPROVE and SEARCH OA observations show little variability among June, July and August, despite large MMV of sulfate in early 2000s (Figure 2A). We find similar behavior from another observation network, CSN. The discontinuity in OA trend in the CSN network is due to different protocols applied (Figure S2). Within sites using the same protocol, there are no systematic monthly differences,
292 293 294 295 296 297	A main constraint comes from the MMV of OA in the southeast US. IMPROVE and SEARCH OA observations show little variability among June, July and August, despite large MMV of sulfate in early 2000s (Figure 2A). We find similar behavior from another observation network, CSN. The discontinuity in OA trend in the CSN network is due to different protocols applied (Figure S2). Within sites using the same protocol, there are no systematic monthly differences, which agrees with IMPROVE and SEARCH. In contrast, modeled OA displays large MMV
<ol> <li>292</li> <li>293</li> <li>294</li> <li>295</li> <li>296</li> <li>297</li> <li>298</li> </ol>	A main constraint comes from the MMV of OA in the southeast US. IMPROVE and SEARCH OA observations show little variability among June, July and August, despite large MMV of sulfate in early 2000s (Figure 2A). We find similar behavior from another observation network, CSN. The discontinuity in OA trend in the CSN network is due to different protocols applied (Figure S2). Within sites using the same protocol, there are no systematic monthly differences, which agrees with IMPROVE and SEARCH. In contrast, modeled OA displays large MMV between June, July and August from 2000 to 2008, where OA in July and August is 1~3 times of
<ol> <li>292</li> <li>293</li> <li>294</li> <li>295</li> <li>296</li> <li>297</li> <li>298</li> <li>299</li> </ol>	A main constraint comes from the MMV of OA in the southeast US. IMPROVE and SEARCH OA observations show little variability among June, July and August, despite large MMV of sulfate in early 2000s (Figure 2A). We find similar behavior from another observation network, CSN. The discontinuity in OA trend in the CSN network is due to different protocols applied (Figure S2). Within sites using the same protocol, there are no systematic monthly differences, which agrees with IMPROVE and SEARCH. In contrast, modeled OA displays large MMV between June, July and August from 2000 to 2008, where OA in July and August is 1~3 times of June values (Figure 2A). Such large MMV is dominated by aqueous SOA, especially from the

2 higher than the observed total OA (Figure 2). The other components including POA and dry 301

302	SOA (including terpene-SOA and Anthropogenic SOA) formed through partitioning together
303	have low concentrations and small MMV. The default model well captures the variability of
304	observed sulfate (Figure 2A), with an average of 3.8 $\mu g/m^3$ and a trend of -6.9%/year, as
305	compared to -6.9%/year (average concentration 4.2 $\mu g/m^3)$ from IMPROVE and -6.7%/year
306	(average concentration 4.3 $\mu$ g/m <sup>3</sup> ) from SEARCH.
307	
308	The large MMV in the model suggests a much stronger modeled OA dependence on sulfate than
309	observations. In 2000-2004, changes in modeled sulfate from June to July and/or August
310	correspond to large MMV of modeled OA mass. In contrast, little MMV is found in observed
311	OA mass during the same months despite large MMV in observed sulfate (Figure 2A). From a
312	linear regression analysis using all monthly data in 2000-2013, the OA-to-sulfate regression
313	slope is <i>m</i> =0.29 ( <i>r</i> <sup>2</sup> =0.25) from IMPROVE, <i>m</i> =0.51 ( <i>r</i> <sup>2</sup> =0.43) from SEARCH, and <i>m</i> =1.87
314	$(r^2=0.57)$ from the default model, even though the default model well captures the magnitude,
315	trend, and monthly variability of observed sulfate. In summary, simulated total OA mass in the
316	standard GEOS-Chem model, dominated by IEPOX-SOA, has a steeper decreasing trend from
317	2000 to 2013 than the observations, and has a large MMV indicating strong dependence on
318	sulfate.
319	

#### 320 3.2 What controls the modeled IEPOX-SOA variability?

The strong dependence of IEPOX-SOA on sulfate is well-established by laboratory and field
work: wet sulfate particles provide the surface and volume of liquid media for IEPOX reactive
uptake (Budisulistiorini et al., 2017; Eddingsaas et al., 2010; Riva et al., 2016; Xu et al., 2015b,
2016), and serve as nucleophiles for nucleophilic addition to form organosulfates (Nguyen et al.,

325	2014; Surratt et al., 2007b). Sulfate (SO <sub>4</sub> <sup>2-</sup> ), together with ammonium (NH <sub>4</sub> <sup>+</sup> ), nitrate (NO <sub>3</sub> <sup>-</sup> ) and
326	other ions, regulates proton (H <sup>+</sup> ) activity $(a_{H+})$ that can catalyze the ring-opening of epoxide
327	group leading to the formation of IEPOX-SOA (Gaston et al., 2014; Pye et al., 2013; Surratt et
328	al., 2007a). However, some recent studies suggest that IEPOX-SOA is not well correlated with
329	aerosol acidity estimated from thermodynamic models (Budisulistiorini et al., 2015; Lin et al.,
330	2013; Xu et al., 2015b), although the lack of direct measurements of aerosol acidity may be a
331	limitation. We use the GEOS-Chem model here to examine the simulated IEPOX-SOA
332	dependence on sulfate, aerosol acidity, and emissions of isoprene which produce IEPOX at high
333	yields under low-NO <sub>x</sub> conditions (Paulot et al., 2009). <u>Temperature impacts the formation of</u>
334	IEPOX-SOA mainly through regulating isoprene emissions but does not influence partitioning as
335	IEPOX-SOA is treated as non-volatile in GEOS-Chem. Therefore, temperature is not examined
336	as another driver in addition to isoprene emissions. We do not treat aerosol water as an
337	independent driver because the dilution effect of aerosol water is implicitly considered in the
338	inorganic sulfate-ammonium-nitrate aerosol volume and acidity calculation, and studies have
339	shown that particle water is not a limiting factor unless the particle is purely dry (Nguyen et al.,
340	2014; Riva et al., 2016; Xu et al., 2015b) which is rare in summertime in the southeast US.
341	
342	We find that the large MMV of OA in the model is mainly driven by sulfate concentrations and
343	aerosol acidity. Figure 3 shows the standardized monthly surface IEPOX-SOA concentration,
344	sulfate concentration, aerosol H <sup>+</sup> activity and isoprene emission from the default model. For each
345	variable, the monthly gridded data has been first averaged over the southeast US. Then, we
346	calculate the one standard deviation of all monthly data (June, July and August data from 2000 to
347	

348	the variables are standardized to be unitless and their variability can be compared directly. Prior
349	to 2008, IEPOX-SOA production is largely enhanced by abundant sulfate (Gaston et al., 2014).
350	Due to this high level of sulfate (about >4 $\mu$ g/m <sup>3</sup> ), the modeled aerosol acidity becomes
351	particularly sensitive to variations in $NH_3$ emissions. The default $NH_3$ emissions from $NEI11v1$
352	suggest no significant long-term trend from 2000 to 2013. In general, ammonium aerosol is
353	strongly correlated with sulfate and has a similar declining trend as sulfate (Silvern et al., 2017).
354	<u>However</u> , the NH <sub>3</sub> emissions in August are about 25% lower than in June and July (Figure S4).
355	As a result, in August before 2008, the aerosol $NH_4^+/SO_4^{2-}$ ratio is smaller (Figure S4) and $a_{H^+}$ is
356	up to 3 times higher than June (Figure 4B), leading to high production of IEPOX-SOA in
357	August. Both sulfate and aerosol acidity appear to be the dominant contributors to MMV of OA
358	during this period. After 2008, IEPOX-SOA formation is substantially suppressed, due to small
359	SO <sub>2</sub> emissions and low modeled aerosol acidity $a_{H^+}$ with small monthly variability. Isoprene
360	emissions also contribute to the month-to-month and interannual OA variability in the model.
361	
362	The multivariate linear regression analysis of IEPOX-SOA quantitatively determines the relative
363	importance of its three drivers in the model. Using all monthly data in 2000-2013, the
364	standardized regression coefficients ( $\beta$ ) associated with $a_{H^+}$ , sulfate aerosol concentration and
365	isoprene emission are $\beta$ =0.50 ( $r^2$ =0.71), $\beta$ =0.39 ( $r^2$ =0.64) and $\beta$ =0.34 ( $r^2$ =0.18), respectively,
366	suggesting that aerosol acidity is the dominant controlling factor. The three variables together
367	explain 88% of the variability of IEPOX-SOA. Their relative importance changes over time
368	(Table S1). Aerosol acidity strongly correlates with IEPOX-SOA in 2005-2008 ( $\beta$ =0.57,
369	$r^2=0.82$ ) but its role becomes much weaker after 2008 ( $\beta=0.27$ , $r^2=0.56$ ). Sulfate aerosol is
370	always the first or second most important driver, especially in 2000-2004 ( $\beta$ =0.46, $r^2$ =0.76).

371	Isoprene emission contributes to the overall interannual variability, for example leading to the
372	relatively low IEPOX-SOA in 2003-2004 and the peaks in 2000, 2006 and 2011 (Figure 3). The
373	high IEPOX-SOA in 2000-2001 and 2005-2007 is a result of high sulfate aerosol, high aerosol
374	acidity due to low NH <sub>3</sub> supply relative to high sulfate, and high isoprene emissions during these
375	periods (Figure 3, Figure 4B).
376	
377	3.3 Narrowing the gap between model and observation
378	3.3.1 Coating
379	Several reasons may lead to the large monthly variations of the modeled OA. The modeled
380	IEPOX-SOA shows a much stronger sensitivity to aerosol acidity than suggested by field
381	observations, which found weak or no correlation between observed IEPOX-SOA and derived
382	aerosol acidity (Budisulistiorini et al., 2015; Lin et al., 2013; Worton et al., 2013; Xu et al.,
383	2015b). Lack of consideration of organic coating effect may provide one possible explanation. In
384	the real atmosphere, inorganic aerosol is generally internally mixed with other organics (Anttila
385	et al., 2006; Murphy et al., 2006). The presence of an organic coating may alter the solubility and
386	diffusion properties at the surface of inorganic particles and diminish further uptake of IEPOX.
387	We implemented a linear coating effect for the IEPOX uptake in a sensitivity simulation 'CT', in
388	which both the magnitude of $\gamma_{EPOX}$ and its sensitivity to acidity have been reduced. Figure 4A
389	shows a schematic illustrating the dependence of the $\gamma_{IEPOX}$ coating effect on acidity $a_{H^+}$ and
390	organic mass fraction ( $\chi_{org}$ ). The original $\gamma_{IEPOX}$ without coating is represented at $\chi_{org}=0$ . The
391	orange line in Figure 4A shows the approximate position of JJA-averaged acidity and organic
392	mass fraction in the <u>'CT'</u> simulation. Adding a coating reduces $\gamma_{IEPOX}$ by almost half, but the
393	impact on the total reactive uptake rate of IEPOX is partially compensated by the corresponding

394increase in particle surface area. The sensitivity of  $\gamma_{IEPOX}$  to acidity has also been reduced395especially during the early 2000s (Figure 4A). The CT simulation reduces the southeast US JJA-396averaged IEPOX-SOA concentrations by  $0.3 \sim 1.8 \ \mu g/m^3$  (Figure 4C).

397

#### 398 3.3.2 NH<sub>3</sub> emissions and aerosol acidity

399 Second, recent studies present contradictory results and explanations on the long-term trend of 400 aerosol acidity in the southeast US (Pye et al., 2020; Silvern et al., 2017; Weber et al., 2016). In 401 this study, we show that the decreasing trend of aerosol acidity from the standard GEOS-Chem 402 model is mainly caused by high acidity in August before 2008, which corresponds to insufficient 403 NH3 emissions in high sulfate environments. The NEI11v1 inventory is used in the default 404 configuration, in which NH<sub>3</sub> emissions in June and July are 30% higher than in August (Figure 405 S4), but not all NH<sub>3</sub> emission inventories agree with such pattern (Paulot et al., 2014). We did a 406 sensitivity test ('CT newNH3') replacing the default US NH3 emissions from NEI11v1 by a new 407 NH<sub>3</sub> emission product derived from CrIS satellite observations, which has higher emissions and 408 smaller MMV among June, July and August (Figure S4). In the 'CT newNH<sub>3</sub>' simulation, the 409 resulting simulated aerosol acidity is substantially changed in 2000-2008 (Figure 4B). The high 410 acidity  $(a_{H+}=0.55\sim0.9 \text{ mol/L})$  in August has been reduced to around 0.2 mol/L and is much 411 closer to June and July values (Figure 3). The results suggest that the fine particles in the 412 southeast US are within a regime where the acidity ( $a_{H+}$  in units of mol/L) is sensitive to NH<sub>3</sub> 413 emissions relative to sulfate concentration, though corresponding pH changes are small (pH 414 within 0.5~1.5, Figure S4). Small changes in NH<sub>3</sub> may lead to large changes in  $a_{H+}$  especially 415 when sulfate concentrations are high, resulting in high month-to-month variability of the IEPOX 416 uptake. After updating the NH3 emissions using the satellite-based estimates, the model

41/	simulates a much more stable trend in aerosol acidity from 2000 to 2015 (Figure 4B), consistent	
418	with recent thermodynamic modeling studies that suggested steady aerosol acidity despite large	
419	reductions in observed sulfate (Pye et al., 2020; Weber et al., 2016).	
420		
421	Due to the high uncertainty associated with the derived NH3 emission product and acidity	
422	calculation (Guo et al., 2015, 2018; Silvern et al., 2017; Song et al., 2018; Tao and Murphy,	
423	2019), we conducted another simulation 'CT_H01' that fix $a_{H^+}$ level at 0.1 mol/L when	
424	calculating IEPOX uptake rate, corresponding to the <u>predicted</u> $a_{H+}$ value (constrained by	Deleted: observed
425	observations) during the 2013 SOAS campaign (Weber et al., 2016). The two simulations,	
426	$CT_newNH_3$ and $CT_H01$ , yield similar long-term trends of IEPOX-SOA in the southeast US	
427	(Figure S5), and they agree better with the long-term surface OA measurements from IMPROVE	
428	and SEARCH than the default model (Figure 4C and 4D). For the SOAS2013 campaign, the	
429	CT_H01 scheme simulates an average IEPOX-SOA concentration of 0.74 $\mu$ g/m <sup>3</sup> , similar to 0.81	
430	$\mu g/m^3$ in the default model, and agrees well with the two independent Aerosol Mass	
431	Spectrometer measurements (0.97 $\mu g/m^3$ from obs_GT and 0.68 $\mu g/m^3$ from obs_CU, see daily	
432	time series in Figure S $\underline{0}$ ). The CT_newNH <sub>3</sub> scheme simulates an average IEPOX-SOA	
433	concentration of 0.34 $\mu$ g/m <sup>3</sup> , lower than the observation and the other models by a factor of >2,	
434	due to both the simplified coating effect and small aerosol $a_{H^+}$ values ( $a_{H^+} < 0.1 \text{ mol/L}$ , Figure	
435	4B). In general, the fixed acidity in the CT_H01 simulation well captures the measured IEPOX-	
436	SOA from the SOAS2013 campaign (Figure S6), and improves the modeled total OA mass	
437	relative to the observations: The modeled long-term decreasing rate of JJA-average OA from	
438	2000 to 2013 has been reduced from 4.9%/year to 3.2%/year, better compared to the IMPROVE	

agal agidity from 2000 to 2012 (Eiguna

1.1

1 .

417

(1D)

20

440 (1.7%/year) and SEARCH (1.9%/year) observations, but is still higher (Figure 4C). The modeled

- 441 MMV of OA have also been greatly reduced (Figure 4D).
- 442

#### 443 **3.3.3 Relationships between OA and sulfate**

444 The formation of aqueous SOA explicitly depends on sulfate aerosol and aerosol acidity which is 445 also impacted by sulfate. The default model, in which a large fraction of simulated total OA mass 446 is from aqueous SOA (mostly IEPOX-SOA), shows a stronger dependence of total OA on sulfate 447 than the observations (Figure 5). The OA-to-sulfate regression slope calculated using monthly 448 OA and sulfate (averaged from all sites beforehand for each network) is m=1.87 for the default simulation, much higher than m=0.29 from IMPROVE and m=0.51 from SEARCH. Such strong 449 450 dependence is clearly demonstrated by the MMV of IEPOX-SOA (Figure 2). Adding the 451 coating effect and fixing  $a_{H+}=0.1$  mol/L substantially reduces the MMV of IEPOX-SOA and the 452 simulated monthly OA-to-sulfate slope (m=1.02). 453 454 Despite the model improvement against the observations in terms of OA and IEPOX-SOA 455 magnitude and long-term relationship with sulfate, the CT H01 scheme needs to be further 456 improved. The rate of OA decreases per year in CT H01 is about 0.8 times higher than the long-457 term observations, with modeled MMV still larger than observations in early 2000s (Figure 4D). 458 Recent studies (Riva et al., 2019) suggested that the IEPOX-SOA production per unit mass of 459 sulfate likely increases with decreasing sulfate due to changes in aerosol properties, such as 460 acidity, morphology, phase state and viscosity, as well as formation of organosulfates, suggesting 461 non-linearity between IEPOX-SOA and sulfate (Riva et al., 2019; Zhang et al., 2019a). Further

462 modeling studies with separated IEPOX-SOA species and detailed aerosol properties are needed463 to achieve a better mechanistic understanding of the dependence of OA on inorganic aerosol.

464

## 465 **4. Summary and Discussion**

466 Significant reduction of SO2 emissions, combined with monthly variations of sulfate and NH3 467 emissions, provide a unique dataset to test the sensitivity of biogenic SOA formation to inorganic 468 species. Observations from two networks (IMPROVE and SEARCH) show a slowly decreasing 469 trend in total OA mass from 2000 to 2013 in the southeast US (-1.7%/year from IMPROVE and -470 1.9%/year from SEARCH), in contrast to a much faster rate of sulfate reduction (-6.9%/year 471 from IMPROVE and -6.7%/year from SEARCH). The standard version of GEOS-Chem model 472 was able to reproduce the long-term trend of sulfate (-6.7%/year), but with a faster decrease of 473 OA (-4.9%/year) and larger interannual variability.

474

475	The MMV of total OA mass during summers provides a novel observational constraint on SOA
476	formation mechanism. Remarkably, we find little MMV of OA from all three surface networks
477	(IMPROVE, SEARCH and CSN) during summer months in 2000-2013, despite larger MMV in
478	sulfate and $NH_3$ emissions. This is in contrast to the standard version of the GEOS-Chem model,
479	which shows a much larger MMV of OA during 2000-2008. Large MMV of OA in the standard
480	model is mainly due to the high sensitivity of modeled IEPOX-SOA to sulfate and aerosol
481	acidity (and $NH_3$ emissions) when sulfate aerosol is abundant. The resulting strong correlation
482	between OA and sulfate also appears to be at odds with long-term observations (Figure 5).
483	Incorporating a coating effect for IEPOX uptake and fixing aerosol acidity, have together

484 improved the model performance in terms of OA trend, variability and the relationship between

485 OA and sulfate, though further improvement is needed.

486

487 There are many uncertainties associated with the calculation of IEPOX-SOA formation. In the 488 default scheme, the Henry's law constant for IEPOX uptake was tuned using measurements from 489 the SOAS2013 campaign and was found to be  $1.7 \times 10^7$  M/atm, 10 times smaller than suggested 490 by Gaston et al. (2014) based on laboratory experiments and about half of the suggested value 491 (3×10<sup>7</sup> M/atm) in some other studies (Budisulistiorini et al., 2017; Nguyen et al., 2014; Pye et 492 al., 2017; Woo and McNeill, 2015; Zhang et al., 2018b). The default simulation agrees well with 493 surface IEPOX-SOA data from SOAS2013 and SEAC4RS 2013 aircraft campaigns (Marais et 494 al., 2016) but overestimates OA magnitude and MMV against long-term observations from 495 IMPROVE and SEARCH. The CT newNH3 simulation reproduces the long-term OA trend but 496 underestimates IEPOX-SOA by a factor of 2 against SOAS 2013. The coating effect may be 497 stronger than used here, as Gaston et al. (2014) used a low viscosity organic material in the 498 experiments. The NH<sub>3</sub> emissions (which are critical for the calculation of aerosol acidity) are 499 highly uncertain (Dammers et al., 2019), and the acidity calculation is further complicated by 500 non-volatile cations (Guo et al., 2018) and meteorological conditions (Guo et al., 2015; Tao and 501 Murphy, 2019). Uncertainties are also associated with the volatility of IEPOX-SOA. Some 502 studies suggested a large fraction of IEPOX-SOA compounds (e.g. 2-methyltetrol) are semi-503 volatile and can re-evaporate back into gas-phase (Ambro et al., 2019; Isaacman-VanWertz et 504 al., 2016), while other studies suggest IEPOX-SOA products are mostly nonvolatile or low 505 volatility (Hu et al., 2016; Lopez-Hilfiker et al., 2016). As multiple parameters may be tuned in 506 the model to fit observations, further laboratory, field and modeling studies are needed to

507	integrate Henry's law constant, IEPOX-SOA yields, volatility, coating effect and acidity	
508	dependence for a better mechanistic understanding. The CT_H01 scheme lacks mechanical	
509	representation of detailed aerosol properties like phase state, acidity, viscosity and morphology,	
510	but reasonably captures both the OA and IEPOX-SOA magnitude (compared to both the three	
511	filter measurement networks and the SOAS2013 campaign), long-term variability and	
512	relationship with sulfate (Figure 4, 5, $S_{0}$ ), therefore may serve as a simplified representation for	
513	climate models. Simulations in this study are conducted at a horizontal resolution of 2°×2.5°,	
514	which is comparable to most global climate models. However, as shown by Yu et al. (2016),	
515	from coarse to fine horizontal resolution, there will be a shift from low-NO <sub>x</sub> to high-NO <sub>x</sub>	
516	pathway for isoprene oxidation. Therefore, using a fine resolution may reduce the production of	
517	IEPOX and IEPOX-SOA, which needs further investigation. For all kinds of models, long-term	
518	filter-based measurements, especially intraseasonal MMV, are important observational	
519	constraints that should be considered in model development.	
520		
521	Even with our improved model, the rate of OA decrease per year is still 0.8 times higher the	
522	long-term observations, and still shows a higher MMV than observations particularly in early	
523	2000s (Figure 4D). Such discrepancies may suggest a more important role of SOA pathways that	
524	are less dependent on inorganic aerosol, such as terpene-SOA formed by reversible gas-aerosol	
525	partitioning. Terpene-SOA is included in GEOS-Chem (yellow color in Figure 1), and	
526	contributes to 8~24% of total OA, which might be underestimated compared to recent field	
527	studies. Xu et al. (2015a) finds a large MMV in IEPOX-SOA, but the less-oxidized oxygenated	Deleted: Despite a large MMV in IEPOX-SOA,
528	OA (LO-OOA, an indicator for freshly-formed monoterpene SOA) and the more-oxidized	
529	oxygenated OA (MO-OOA, also likely from biogenic sources) have little MMV in summer	

- 531 months, and they contribute to more than 50% of total OA mass in the southeast US (Xu et al.,
- 532 2018). The important role of monoterpenes SOA is also confirmed by molecular level
- 533 characterization of organic aerosols (Zhang et al., 2018a). Other pathways may contribute to
- 534 SOA to some extent and may add to the predicted SOA formed by partitioning, including
- 535 biogenic SOA from auto-oxidation (Bianchi et al., 2019; Pye et al., 2019), in-cloud SOA
- 536 formation that may be less dependent on acidity than aqueous SOA (Tsui et al., 2019), a small
- 537 but underestimated contribution of anthropogenic SOA (Schroder et al., 2018; Shah et al., 2019),
- and other possible mechanisms (Schwantes et al., 2019). Further quantifying the relative
- 539 importance of the different pathways will allow a more accurate quantification of the
- 540 anthropogenic influence on biogenic SOA and the associated radiative forcing.
- 541

542

#### 543 Data availability

- 544 The observational datasets from long-term filter measurement networks IMPROVE and CSN are
- 545 available at http://views.cira.colostate.edu/fed/QueryWizard/Default.aspx. The SEARCH
- 546 observational datasets are available by contacting E. Edgerton. The model code and modeling
- 547 results are available by contacting Y. Zheng and J. Mao.

#### 549 Acknowledgement

548

557

- 550 YZ and JM acknowledge funding from NOAA NA18OAR4310114. HC and DKJ recognize
- 551 support from NASA 80NSSC18K0689. WH and JLJ acknowledge funding from NSF AGS-
- 552 1822664. EAM acknowledges funding from NERC/EPSRC (award number EP/R513465/1). The
- authors acknowledge the Electric Power Research Institute (EPRI) and Southern Company for
- 554 support of the SEARCH network and Atmospheric Research & Analysis. IMPROVE and CSN
- data are accessed from the Federal Land Manger Environmental Database. YZ thanks helpful
- 556 discussions with Arlene M. Fiore, Róisín Commane and V. Faye McNeill.

#### 558 Author Contributions:

- 559 Y.Z. and J.M. designed the research, performed the simulations and conducted the analysis.
- 560 J.A.T. and E.M.M. provided guidance on aerosol coating parameterization. H.C. and D.K.H.
- 561 provided the CrIS-derived NH<sub>3</sub> emission. N.L.N, W.H. and J.L.J. provided data from the
- 562 SOAS2013 field campaign. E.E. provided data from the SEARCH network. Y.Z. wrote the paper
- with all coauthors providing input.

#### 565 Competing interests

566 The authors declare no competing interests.

#### 568 Additional information

569 Correspondence and material requests should be addressed to Y. Zheng and J. Mao.

570

567

#### 571 Reference:

- 572 Ambro, E. L., Schobesberger, S., Gaston, C. J., Lopez-Hilfiker, F. D., Lee, B. H., Liu, J.,
- 573 Zelenyuk, A., Bell, D., Cappa, C. D., Helgestad, T., Li, Z., Guenther, A., Wang, J., Wise, M.,
- 574 Caylor, R., Surratt, J. D., Riedel, T., Hyttinen, N., Salo, V.-T., Hasan, G., Kurtén, T., Shilling, J.
- 575 E. and Thornton, J. A.: Chamber-based insights into the factors controlling IEPOX SOA yield,
- 576 composition, and volatility, Atmos. Chem. Phys., 19, 11253–11265, doi:10.5194/acp-2019-271, 577 2019.
- 578 Anttila, T., Kiendler-Scharr, A., Tillmann, R. and Mentel, T. F.: On the reactive uptake of
- 579 gaseous compounds by organic-coated aqueous aerosols: Theoretical analysis and application to
- 580 the heterogeneous hydrolysis of N2O5, J. Phys. Chem. A, 110(35), 10435–10443,
- 581 doi:10.1021/jp062403c, 2006.
- 582 Attwood, A. R., Washenfelder, R. A., Brock, C. A., Hu, W., Baumann, K., Campuzano-Jost, P.,
- 583 Day, D. A., Edgerton, E. S., Murphy, D. M., Palm, B. B., McComiskey, A., Wagner, N. L., De
- 584 Sá, S. S., Ortega, A., Martin, S. T., Jimenez, J. L. and Brown, S. S.: Trends in sulfate and organic
- 585 aerosol mass in the Southeast U.S.: Impact on aerosol optical depth and radiative forcing,
- 586 Geophys. Res. Lett., 41, 7701–7709, doi:10.1002/2014GL061669, 2014.
- 587 Bey, I., Jacob, D. J., Yantosca, R. M., Logan, J. A., Field, B. D., Fiore, A. M., Li, Q.-B., Liu, H.-
- 588 Y., Mickley, L. J. and Schultz, M. G.: Global Modeling of Tropospheric Chemistry with
- Assimilated Meteorology: Model Description and Evaluation, J. Geophys. Res., 106, 73–95,
   doi:10.1029/2001JD000807, 2001.
- 591 Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crounse, J.
- 592 D., Wennberg, P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop,
- 593 D. R., Thornton, J. A., Donahue, N., Kjaergaard, H. G. and Ehn, M.: Highly Oxygenated Organic
- 594 Molecules (HOM) from Gas-Phase Autoxidation Involving Peroxy Radicals: A Key Contributor
- to Atmospheric Aerosol, Chem. Rev., 119(6), 3472–3509, doi:10.1021/acs.chemrev.8b00395,
   2019.
- Blanchard, C. L., Hidy, G. M., Shaw, S., Baumann, K. and Edgerton, E. S.: Effects of emission
   reductions on organic aerosol in the southeastern United States, Atmos. Chem. Phys., 16, 215–
- 599 238, doi:10.5194/acp-16-215-2016, 2016.
- Budisulistiorini, S. H., Li, X., Bairai, S. T., Renfro, J., Liu, Y., Liu, Y. J., McKinney, K. A.,
- Martin, S. T., McNeill, V. F., Pye, H. O. T., Nenes, A., Neff, M. E., Stone, E. A., Mueller, S.,
- 602 Knote, C., Shaw, S. L., Zhang, Z., Gold, A. and Surratt, J. D.: Examining the effects of
- anthropogenic emissions on isoprene-derived secondary organic aerosol formation during the
- 2013 Southern Oxidant and Aerosol Study (SOAS) at the Look Rock, Tennessee ground site,
  Atmos. Chem. Phys., 15(15), 8871–8888, doi:10.5194/acp-15-8871-2015, 2015.
- Budisulistiorini, S. H., Baumann, K., Edgerton, E. S., Bairai, S. T., Mueller, S., Shaw, S. L.,
- 607 Knipping, E. M., Gold, A. and Surratt, J. D.: Seasonal characterization of submicron aerosol
- 608 chemical composition and organic aerosol sources in the southeastern United States: Atlanta,
- 609 Georgia, and Look Rock, Tennessee, Atmos. Chem. Phys., 16(8), 5171–5189, doi:10.5194/acp-
- 610 16-5171-2016, 2016.
- 611 Budisulistiorini, S. H., Nenes, A., Carlton, A. G., Surratt, J. D., McNeill, V. F. and Pye, H. O. T.:
- Simulating Aqueous-Phase Isoprene-Epoxydiol (IEPOX) Secondary Organic Aerosol Production
   during the 2013 Southern Oxidant and Aerosol Study (SOAS), Environ. Sci. Technol., 51(9),
- 614 5026–5034, doi:10.1021/acs.est.6b05750, 2017.
- 5020 = 5054, 401.1021/acs.est.0005/50, 2017.
- 615 Butler, T., Vermeylen, F., Lehmann, C. M., Likens, G. E. and Puchalski, M.: Increasing
- 616 ammonia concentration trends in large regions of the USA derived from the NADP/AMoN

- 617 network, Atmos. Environ., 146(3), 132–140, doi:10.1016/j.atmosenv.2016.06.033, 2016.
- 618 Cao, H., Henze, D. K., Shephard, M. W., Dammers, E., Cady-Pereira, K., Alvarado, M. J.,
- 619 Lonsdale, C. R., Luo, G., Yu, F., Zhu, L., Danielson, C. G. and Edgerton, E. S.: Inverse
- 620 modeling of NH3 sources using CrIS remote sensing measurements, Environ. Res. Lett., in
- 621 press, doi:https://doi.org/10.1088/1748-9326/abb5cc, 2020.
- 622 Carlton, A. G., Pinder, R. W., Bhave, P. V. and Pouliot, G. A.: To what extent can biogenic SOA
- 623 be controlled?, Environ. Sci. Technol., 44(9), 3376–3380, doi:10.1021/es903506b, 2010.
- 624 Dammers, E., McLinden, C. A., Griffin, D., Shephard, M. W., Van Der Graaf, S., Lutsch, E.,
- 625 Schaap, M., Gainairu-Matz, Y., Fioletov, V., Van Damme, M., Whitburn, S., Clarisse, L., Cady-
- 626 Pereira, K., Clerbaux, C., Francois Coheur, P. and Erisman, J. W.: NH3 emissions from large
- 627 point sources derived from CrIS and IASI satellite observations, Atmos. Chem. Phys., 19(19),
- 628 12261–12293, doi:10.5194/acp-19-12261-2019, 2019.
- 629 Donahue, N. M., Robinson, a. L., Stanier, C. O. and Pandis, S. N.: Coupled partitioning,
- dilution, and chemical aging of semivolatile organics, Environ. Sci. Technol., 40(8), 2635–2643,
  doi:10.1021/es052297c, 2006.
- 632 Eddingsaas, N. C., Vandervelde, D. G. and Wennberg, P. O.: Kinetics and products of the acid-
- 633 catalyzed ring-opening of atmospherically relevant butyl epoxy alcohols, J. Phys. Chem. A, 114,
- 634 8106–8113, doi:10.1021/jp103907c, 2010.
- 635 Edgerton, E. S., Hartsell, B. E., Saylor, R. D., Jansen, J. J., Hansen, D. A. and Hidy, G. M.: The
- 636 southeastern aerosol research and characterization study: Part II. Filter-based measurements of
- 637 fine and coarse particulate matter mass and composition, J. Air Waste Manag. Assoc., 55(10),
- 638 1527–1542, doi:10.1080/10473289.2005.10464744, 2005.
- EPA: The benefits and costs of the clean air act from 1990-2020, in Better Air: Benefits andCosts of the Clean Air Act., 2011.
- 040 Cosis of the Clean All Act., 2011.
- 641 Ervens, B., Turpin, B. J. and Weber, R. J.: Secondary organic aerosol formation in cloud droplets
- 642 and aqueous particles (aqSOA): A review of laboratory, field and model studies, Atmos. Chem.
- 643 Phys., 11(21), 11069–11102, doi:10.5194/acp-11-11069-2011, 2011.
- 644 Fisher, J. A., Jacob, D. J., Travis, K. R., Kim, P. S., Marais, E. A., Miller, C. C., Yu, K., Zhu, L.,
- 645 Yantosca, R. M., Sulprizio, M. P., Mao, J., Wennberg, P. O., Crounse, J. D., Teng, A. P.,
- 646 Nguyen, T. B., Clair, J. M. S., Cohen, R. C., Romer, P., Nault, B. A., Wooldridge, P. J., Jimenez,
- 647 J. L., Campuzano-Jost, P., Day, D. A., Hu, W., Shepson, P. B., Xiong, F., Blake, D. R.,
- 648 Goldstein, A. H., Misztal, P. K., Hanisco, T. F., Wolfe, G. M., Ryerson, T. B., Wisthaler, A. and
- 649 Mikoviny, T.: Organic nitrate chemistry and its implications for nitrogen budgets in an isoprene-
- and monoterpene-rich atmosphere: Constraints from aircraft (SEAC4RS) and ground-based
- 651 (SOAS) observations in the Southeast US, Atmos. Chem. Phys., 16(9), 5969–5991,
- 652 doi:10.5194/acp-16-5969-2016, 2016.
- 653 Fountoukis, C. and Nenes, A.: ISORROPIA II: a computationally efficient thermodynamic
- 654 equilibrium model for K+-Ca2+-Mg2+-NH4+-Na+-SO42--NO3--Cl--H2O aerosols, Atmos.
- 655 Chem. Phys, 7, 4639–4659, doi:10.5194/acp-7-4639-2007, 2007.
- 656 Fu, T. M., Jacob, D. J., Wittrock, F., Burrows, J. P., Vrekoussis, M. and Henze, D. K.: Global
- budgets of atmospheric glyoxal and methylglyoxal, and implications for formation of secondary
- organic aerosols, J. Geophys. Res. Atmos., doi:10.1029/2007JD009505, 2008.
- 659 Gaston, C. J., Riedel, T. P., Zhang, Z., Gold, A., Surratt, J. D. and Thornton, J. A.: Reactive
- uptake of an isoprene-derived epoxydiol to submicron aerosol particles, Environ. Sci. Technol.,
  48(19), 11178–11186, doi:10.1021/es5034266, 2014.
- 662 Guenther, A. B., Karl, T., Harley, P., Wiedinmyer, C., Palmer, P. I. and Geron, C.: Estimates of

- 663 global terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols
- from Nature), Atmos. Chem. Phys., 6, 3181–3210, doi:doi:10.5194/acp-6-3181-2006, 2006.
- 665 Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L. K. and
- 666 Wang, X.: The model of emissions of gases and aerosols from nature version 2.1 (MEGAN2.1):
- 667 An extended and updated framework for modeling biogenic emissions, Geosci. Model Dev., 5, 668 1471–1492, doi:10.5194/gmd-5-1471-2012, 2012.
- 669 Guo, H., Xu, L., Bougiatioti, A., Cerully, K. M., Capps, S. L., Hite, J. R., Carlton, A. G., Lee, S.
- 670 H., Bergin, M. H., Ng, N. L., Nenes, A. and Weber, R. J.: Fine-particle water and pH in the
- 671 southeastern United States, Atmos. Chem. Phys., 15(9), 5211–5228, doi:10.5194/acp-15-5211-672 2015, 2015.
- 673 Guo, H., Nenes, A. and Weber, R. J.: The underappreciated role of nonvolatile cations in aerosol
- 674 ammonium-sulfate molar ratios, Atmos. Chem. Phys., 18(23), 17307–17323, doi:10.5194/acp-675 18-17307-2018, 2018.
- 676 Hayes, P. L., Carlton, A. G., Baker, K. R., Ahmadov, R., Washenfelder, R. A., Alvarez, S.,
- 677 Rappenglück, B., Gilman, J. B., Kuster, W. C., De Gouw, J. A., Zotter, P., Prévôt, A. S. H.,
- 678 Szidat, S., Kleindienst, T. E., Offenberg, J. H., Ma, P. K. and Jimenez, J. L.: Modeling the
- 679 formation and aging of secondary organic aerosols in Los Angeles during CalNex 2010, Atmos.
- 680 Chem. Phys., 15(10), 5773–5801, doi:10.5194/acp-15-5773-2015, 2015.
- 681 Henze, D. K., Hakami, A. and Seinfeld, J. H.: Development of the adjoint of GEOS-Chem,
- 682 Atmos. Chem. Phys., 7(9), 2413–2433, doi:10.5194/acp-7-2413-2007, 2007.
- 683 Hidy, G. M., Blanchard, C. L., Baumann, K., Edgerton, E., Tanenbaum, S., Shaw, S., Knipping,
- E., Tombach, I., Jansen, J. and Walters, J.: Chemical climatology of the southeastern United
- 685 States, 1999-2013, Atmos. Chem. Phys., 14(21), 11893–11914, doi:10.5194/acp-14-11893-2014, 686 2014.
- 687 Hodshire, A. L., Akherati, A., Alvarado, M. J., Brown-Steiner, B., Jathar, S. H., Jimenez, J. L.,
- 688 Kreidenweis, S. M., Lonsdale, C. R., Onasch, T. B., Ortega, A. M. and Pierce, J. R.: Aging
- 689 Effects on Biomass Burning Aerosol Mass and Composition: A Critical Review of Field and
- Laboratory Studies, Environ. Sci. Technol., 53(17), 10007–10022, doi:10.1021/acs.est.9b02588,
  2019.
- Hu, W., Palm, B. B., Day, D. A., Campuzano-Jost, P., Krechmer, J. E., Peng, Z., De Sa Suzane,
- 693 S., Martin, S. T., Alexander, M. L., Baumann, K., Hacker, L., Kiendler-Scharr, A., Koss, A. R.,
- 694 De Gouw, J. A., Goldstein, A. H., Seco, R., Sjostedt, S. J., Park, J. H., Guenther, A. B., Kim, S.,
- 695 Canonaco, F., Prévôt, A. S. H., Brune, W. H. and Jimenez, J. L.: Volatility and lifetime against
- 696 OH heterogeneous reaction of ambient isoprene-epoxydiols-derived secondary organic aerosol
- 697 (IEPOX-SOA), Atmos. Chem. Phys., 16(18), 11563–11580, doi:10.5194/acp-16-11563-2016,
- 698 2016.
- 699 Hu, W. W., Campuzano-Jost, P., Palm, B. B., Day, D. A., Ortega, A. M., Hayes, P. L.,
- 700 Krechmer, J. E., Chen, Q., Kuwata, M., Liu, Y. J., De Sá, S. S., McKinney, K., Martin, S. T., Hu,
- 701 M., Budisulistiorini, S. H., Riva, M., Surratt, J. D., St. Clair, J. M., Isaacman-Van Wertz, G.,
- 702 Yee, L. D., Goldstein, A. H., Carbone, S., Brito, J., Artaxo, P., De Gouw, J. A., Koss, A.,
- 703 Wisthaler, A., Mikoviny, T., Karl, T., Kaser, L., Jud, W., Hansel, A., Docherty, K. S., Alexander,
- 704 M. L., Robinson, N. H., Coe, H., Allan, J. D., Canagaratna, M. R., Paulot, F. and Jimenez, J. L.:
- 705 Characterization of a real-time tracer for isoprene epoxydiols-derived secondary organic aerosol
- 706 (IEPOX-SOA) from aerosol mass spectrometer measurements, Atmos. Chem. Phys., 15(20),
- 707 11807–11833, doi:10.5194/acp-15-11807-2015, 2015.
- 708 Isaacman-VanWertz, G., Yee, L. D., Kreisberg, N. M., Wernis, R., Moss, J. A., Hering, S. V., De

- 709 Sá, S. S., Martin, S. T., Alexander, M. L., Palm, B. B., Hu, W., Campuzano-Jost, P., Day, D. A.,
- 710 Jimenez, J. L., Riva, M., Surratt, J. D., Viegas, J., Manzi, A., Edgerton, E., Baumann, K., Souza,
- 711 R., Artaxo, P. and Goldstein, A. H.: Ambient Gas-Particle Partitioning of Tracers for Biogenic
- 712 Oxidation, Environ. Sci. Technol., 50(18), 9952–9962, doi:10.1021/acs.est.6b01674, 2016.
- Janssens-Maenhout, G., Crippa, M., Guizzardi, D., Dentener, F., Muntean, M., Pouliot, G.,
   Keating, T., Zhang, O., Kurokawa, J., Wankmüller, R., Denier Van Der Gon, H., Kuenen, J.
- Keating, T., Zhang, Q., Kurokawa, J., Wankmüller, R., Denier Van Der Gon, H., Kuenen, J. J.
  P., Klimont, Z., Frost, G., Darras, S., Koffi, B. and Li, M.: HTAP-v2.2: A mosaic of regional and
- P., Klimont, Z., Frost, G., Darras, S., Koffi, B. and Li, M.: HTAP-v2.2: A mosaic of regional
   global emission grid maps for 2008 and 2010 to study hemispheric transport of air pollution,
- 717 Atmos. Chem. Phys., 15(19), 11411–11432, doi:10.5194/acp-15-11411-2015, 2015.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H.,
- 719 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M.,
- 720 Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin,
- 721 C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M.,
- Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J. A.,
- 723 Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F.,
- 724 Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A.,
- 725 Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J.
- 726 R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C.,
- 727 Middlebrook, A. M., Kolb, C. E., Baltensperger, U. and Worsnop, D. R.: Evolution of organic
- aerosols in the atmosphere, Science (80-.)., 326(5959), 1525–1529,
- 729 doi:10.1126/science.1180353, 2009.
- 730 Jo, D. S., Hodzic, A., Emmons, L. K., Marais, E. A., Peng, Z., Nault, B. A., Hu, W.,
- 731 Campuzano-Jost, P. and Jimenez, J. L .: A simplified parameterization of isoprene-epoxydiol-
- 732 derived secondary organic aerosol (IEPOX-SOA) for global chemistry and climate models: A
- 733 case study with GEOS-Chem v11-02-rc, Geosci. Model Dev., 12(7), 2983–3000,
- 734 doi:10.5194/gmd-12-2983-2019, 2019.
- 735 Kiendler-Scharr, A., Mensah, A. A., Friese, E., Topping, D., Nemitz, E., Prevot, A. S. H., Äijälä,
- 736 M., Allan, J., Canonaco, F., Canagaratna, M., Carbone, S., Crippa, M., Dall Osto, M., Day, D.
- 737 A., De Carlo, P., Di Marco, C. F., Elbern, H., Eriksson, A., Freney, E., Hao, L., Herrmann, H.,
- 738 Hildebrandt, L., Hillamo, R., Jimenez, J. L., Laaksonen, A., McFiggans, G., Mohr, C., O'Dowd,
- 739 C., Otjes, R., Ovadnevaite, J., Pandis, S. N., Poulain, L., Schlag, P., Sellegri, K., Swietlicki, E.,
- 740 Tiitta, P., Vermeulen, A., Wahner, A., Worsnop, D. and Wu, H. C.: Ubiquity of organic nitrates
- from nighttime chemistry in the European submicron aerosol, Geophys. Res. Lett., 43(14), 7735–
  7744, doi:10.1002/2016GL069239, 2016.
- 743 Kim, P. S., Jacob, D. J., Fisher, J. A., Travis, K., Yu, K., Zhu, L., Yantosca, R. M., Sulprizio, M.
- 744 P., Jimenez, J. L., Campuzano-Jost, P., Froyd, K. D., Liao, J., Hair, J. W., Fenn, M. A., Butler, C.
- 745 F., Wagner, N. L., Gordon, T. D., Welti, A., Wennberg, P. O., Crounse, J. D., St. Clair, J. M.,
- 746 Teng, A. P., Millet, D. B., Schwarz, J. P., Markovic, M. Z. and Perring, A. E.: Sources,
- 747 seasonality, and trends of Southeast US aerosol: An integrated analysis of surface, aircraft, and
- ratellite observations with the GEOS-Chem chemical transport model, Atmos. Chem. Phys.
- 749 Discuss., 15(13), 17651–17709, doi:10.5194/acpd-15-17651-2015, 2015.
- 750 Kroll, J. H., Ng, N. L., Murphy, S. M., Flagan, R. C. and Seinfeld, J. H.: Secondary organic
- 751 aerosol formation from isoprene photooxidation under high-NOx conditions, Geophys. Res.
- 752 Lett., 32(x), 1–4, doi:10.1029/2005GL023637, 2005.
- 753 Kroll, J. H., Ng, N. L., Murphy, S. M., Flagan, R. C. and Seinfeld, J. H.: Secondary organic
- aerosol formation from isoprene photooxidation, Environ. Sci. Technol., 40(3), 1869–1877,

- 755 doi:10.1021/es0524301, 2006.
- 756 Li, J., Mao, J., Min, K. E., Washenfelder, R. A., Brown, S. S., Kaiser, J., Keutsch, F. N.,
- 757 Volkamer, R., Wolfe, G. M., Hanisco, T. F., Pollack, I. B., Ryerson, T. B., Graus, M., Gilman, J.
- 758 B., Lerner, B. M., Warneke, C., de Gouw, J. A., Middlebrook, A. M., Liao, J., Welti, A.,
- 759 Henderson, B. H., Faye McNeill, V., Hall, S. R., Ullmann, K., Donner, L. J., Paulot, F. and
- 760 Horowitz, L. W.: Observational constraints on glyoxal production from isoprene oxidation and
- its contribution to organic aerosol over the Southeast United States, J. Geophys. Res., 761
- 762 doi:10.1002/2016JD025331, 2016.
- 763 Li, J., Mao, J., Fiore, A. M., Cohen, R. C., Crounse, J. D., Teng, A. P., Wennberg, P. O., Lee, B.
- 764 H., Lopez-Hilfiker, F. D., Thornton, J. A., Peischl, J., Pollack, I. B., Ryerson, T. B., Veres, P.,
- 765 Roberts, J. M., Neuman, J. A., Nowak, J. B., Wolfe, G. M., Hanisco, T. F., Fried, A., Singh, H.
- 766 B., Dibb, J., Paulot, F. and Horowitz, L. W.: Decadal changes in summertime reactive oxidized
- 767 nitrogen and surface ozone over the Southeast United States, Atmos. Chem. Phys., 18(3), 2341-768 2361, doi:10.5194/acp-18-2341-2018, 2018.
- 769
- Li, Y., Day, D. A., Stark, H., Jimenez, J. L. and Shiraiwa, M.: Predictions of the glass transition 770 temperature and viscosity of organic aerosols from volatility distributions, Atmos. Chem. Phys.,
- 20(13), 8103-8122, doi:10.5194/acp-20-8103-2020, 2020. 771
- 772 Liggio, J., Li, S. M. and McLaren, R.: Reactive uptake of glyoxal by particulate matter, J.
- Geophys. Res. D Atmos., 110(10), 1-13, doi:10.1029/2004JD005113, 2005. 773
- 774 Lin, Y. H., Knipping, E. M., Edgerton, E. S., Shaw, S. L. and Surratt, J. D.: Investigating the
- 775 influences of SO2 and NH3 levels on isoprene-derived secondary organic aerosol formation
- 776 using conditional sampling approaches, Atmos. Chem. Phys., 13(16), 8457-8470,
- 777 doi:10.5194/acp-13-8457-2013, 2013.
- 778 Lopez-Hilfiker, F. D., Mohr, C., D'Ambro, E. L., Lutz, A., Riedel, T. P., Gaston, C. J., Iyer, S.,
- Zhang, Z., Gold, A., Surratt, J. D., Lee, B. H., Kurten, T., Hu, W. W., Jimenez, J., Hallquist, M. 779
- 780 and Thornton, J. A.: Molecular Composition and Volatility of Organic Aerosol in the
- 781 Southeastern U.S.: Implications for IEPOX Derived SOA, Environ. Sci. Technol., 50(5), 2200-782 2209, doi:10.1021/acs.est.5b04769, 2016.
- 783 Malm, W. C., Schichtel, B. A., Hand, J. L. and Collett, J. L.: Concurrent Temporal and Spatial
- 784 Trends in Sulfate and Organic Mass Concentrations Measured in the IMPROVE Monitoring
- 785 Program, J. Geophys, Res. Atmos., 122(19), 10462-10476, doi:10.1002/2017JD026865, 2017.
- 786 Mao, J., Paulot, F., Jacob, D. J., Cohen, R. C., Crounse, J. D., Wennberg, P. O., Keller, C. A.,
- 787 Hudman, R. C., Barkley, M. P. and Horowitz, L. W.: Ozone and organic nitrates over the eastern 788 United States: Sensitivity to isoprene chemistry, J. Geophys. Res. Atmos., 118(19), 11256-
- 789 11268, doi:10.1002/jgrd.50817, 2013.
- 790 Mao, J., Carlton, A., Cohen, R. C., Brune, W. H., Brown, S. S., Wolfe, G. M., Jimenez, J. L.,
- 791 Pye, H. O. T., Lee Ng, N., Xu, L., Faye McNeill, V., Tsigaridis, K., McDonald, B. C., Warneke,
- 792 C., Guenther, A., Alvarado, M. J., De Gouw, J., Mickley, L. J., Leibensperger, E. M., Mathur,
- 793 R., Nolte, C. G., Portmann, R. W., Unger, N., Tosca, M. and Horowitz, L. W.: Southeast
- 794 Atmosphere Studies: Learning from model-observation syntheses, Atmos. Chem. Phys., 18(4),
- 795 2615-2651, doi:10.5194/acp-18-2615-2018, 2018.
- 796 Marais, E. A., Jacob, D. J., Jimenez, J. L., Campuzano-Jost, P., Day, D. A., Hu, W., Krechmer,
- 797 J., Zhu, L., Kim, P. S., Miller, C. C., Fisher, J. A., Travis, K., Yu, K., Hanisco, T. F., Wolfe, G.
- 798 M., Arkinson, H. L., Pve, H. O. T., Froyd, K. D., Liao, J. and McNeill, V. F.: Aqueous-phase
- 799 mechanism for secondary organic aerosol formation from isoprene: Application to the Southeast
- 800 United States and co-benefit of SO2 emission controls, Atmos. Chem. Phys., 16, 1603-1618,

- 801 doi:10.5194/acp-16-1603-2016, 2016.
- 802 Marais, E. A., Jacob, D. J., Turner, J. R. and Mickley, L. J.: Evidence of 1991–2013 decrease of
- 803 biogenic secondary organic aerosol in response to SO2 emission controls, Environ. Sci.
- 804 Technol., 12, doi:https://doi.org/10.1088/1748-9326/aa69c8, 2017.
- 805 McNeill, V. F., Woo, J. L., Kim, D. D., Schwier, A. N., Wannell, N. J., Sumner, A. J. and
- 806 Barakat, J. M.: Aqueous-phase secondary organic aerosol and organosulfate formation in
- atmospheric aerosols: A modeling study, Environ. Sci. Technol., 46(15), 8075–8081,
- 808 doi:10.1021/es3002986, 2012.
- 809 Murphy, D. M., Cziczo, D. J., Froyd, K. D., Hudson, P. K., Matthew, B. M., Middlebrook, A.
- 810 M., Peltier, R. E., Sullivan, A., Thomson, D. S. and Weber, R. J.: Single-peptide mass
- 811 spectrometry of tropospheric aerosol particles, J. Geophys. Res. Atmos., 111(23), 1–15,
- 812 doi:10.1029/2006JD007340, 2006.
- 813 Ng, N. L., Brown, S. S., Archibald, A. T., Atlas, E., Cohen, R. C., Crowley, J. N., Day, D. A.,
- 814 Donahue, N. M., Fry, J. L., Fuchs, H., Griffin, R. J., Guzman, M. I., Herrmann, H., Hodzic, A.,
- 815 Iinuma, Y., Kiendler-Scharr, A., Lee, B. H., Luecken, D. J., Mao, J., McLaren, R., Mutzel, A.,
- 816 Osthoff, H. D., Ouyang, B., Picquet-Varrault, B., Platt, U., Pye, H. O. T., Rudich, Y., Schwantes,
- 817 R. H., Shiraiwa, M., Stutz, J., Thornton, J. A., Tilgner, A., Williams, B. J. and Zaveri, R. A.:
- 818 Nitrate radicals and biogenic volatile organic compounds: Oxidation, mechanisms, and organic
- 819 aerosol, Atmos. Chem. Phys., 17(3), 2103–2162, doi:10.5194/acp-17-2103-2017, 2017.
- 820 Nguyen, T. B., Coggon, M. M., Bates, K. H., Zhang, X., Schwantes, R. H., Schilling, K. A.,
- 821 Loza, C. L., Flagan, R. C., Wennberg, P. O. and Seinfeld, J. H.: Organic aerosol formation from
- the reactive uptake of isoprene epoxydiols (IEPOX) onto non-acidified inorganic seeds, Atmos.
  Chem. Phys., 14(7), 3497–3510, doi:10.5194/acp-14-3497-2014, 2014.
- Pai, S., Heald, C., Pierce, J., Farina, S., Marais, E., Jimenez, J., Campuzano-Jost, P., Nault, B.,
- Middlebrook, A., Coe, H., Shilling, J., Bahreini, R., Dingle, J. and Vu, K.: An evaluation of
- global organic aerosol schemes using airborne observations, Atmos. Chem. Phys., (20), 2637–
- global organic deleter belenes asing anothe observations, rands: chem. Phys., (20)
   2665, doi:10.5194/acp-2019-331, 2020.
- Pankow, J. F.: An Absorption-Model of the Gas Aerosol Partitioning Involved in the Formation
  of Secondary Organic Aerosol, Atmos. Environ., 28(2), 189–193,
- 830 doi:10.1016/j.atmosenv.2007.10.060, 1994.
- 831 Paulot, F., Crounse, J. D., Kjaergaard, H. G., Kürten, A., Clair, J. M. S., Seinfeld, J. H. and
- 832 Wennberg, P. O.: Unexpected Epoxide Formation in the Gas-Phase Photooxidation of Isoprene,
- 833 Science (80-. )., 325, 730–734, doi:10.1126/science.1174251, 2009.
- 834 Paulot, F., Jacob, D. J., Pinder, R. W., Bash, J. O., Travis, K. R. and Henze, D. K.: Ammonia
- 835 emissions in the United States, European Union, and China derived by high-resolution inversion
- 836 of ammonium wet deposition data: Interpretation with a new agricultural emissions inventory
- 837 (MASAGE\_NH3), J. Geophys. Res. Atmos., 119, 4343–4364, doi:doi:10.1002/2013JD021130,
   838 2014.
- 839 Presto, A. a., Huff Hartz, K. E. and Donahue, N. M.: Secondary organic aerosol production from
- terpene ozonolysis. 2. Effect of NOx concentration, Environ. Sci. Technol., 39(18), 7046–7054,
  doi:10.1021/es050400s, 2005.
- 041 doi.10.1021/cs050400s, 200.
- 842 Pye, H. O. T., Liao, H., Wu, S., Mickley, L. J., Jacob, D. J., Henze, D. J. and Seinfeld, J. H.:
- 843 Effect of changes in climate and emissions on future sulfate-nitrate-ammonium aerosol levels in
- the United States, J. Geophys. Res. Atmos., 114(1), 1–18, doi:10.1029/2008JD010701, 2009.
- 845 Pye, H. O. T., Chan, a. W. H., Barkley, M. P. and Seinfeld, J. H.: Global modeling of organic
- aerosol: The importance of reactive nitrogen (NOx and NO3), Atmos. Chem. Phys., 10, 11261-

- 847 11276, doi:10.5194/acp-10-11261-2010, 2010.
- 848 Pye, H. O. T., Pinder, R. W., Piletic, I. R., Xie, Y., Capps, S. L., Lin, Y. H., Surratt, J. D., Zhang,
- 849 Z., Gold, A., Luecken, D. J., Hutzell, W. T., Jaoui, M., Offenberg, J. H., Kleindienst, T. E.,
- 850 Lewandowski, M. and Edney, E. O.: Epoxide pathways improve model predictions of isoprene
- 851 markers and reveal key role of acidity in aerosol formation, Environ. Sci. Technol., 47(19),
- 852 11056–11064, doi:10.1021/es402106h, 2013.
- 853 Pye, H. O. T., Murphy, B. N., Xu, L., Ng, N. L., Carlton, A. G., Guo, H., Weber, R., Vasilakos,
- 854 P., Wyat Appel, K., Hapsari Budisulistiorini, S., Surratt, J. D., Nenes, A., Hu, W., Jimenez, J. L.,
- 855 Isaacman-Vanwertz, G., Misztal, P. K. and Goldstein, A. H.: On the implications of aerosol
- 856 liquid water and phase separation for organic aerosol mass, Atmos. Chem. Phys., 17(1), 343–
- 857 369, doi:10.5194/acp-17-343-2017, 2017.
- 858 Pye, H. O. T., D'Ambro, E. L., Lee, B. H., Schobesberger, S., Takeuchi, M., Zhao, Y., Lopez-
- 859 Hilfiker, F., Liu, J., Shilling, J. E., Xing, J., Mathur, R., Middlebrook, A. M., Liao, J., Welti, A.,
- 860 Graus, M., Warneke, C., de Gouw, J. A., Holloway, J. S., Ryerson, T. B., Pollack, I. B. and
- Thornton, J. A.: Anthropogenic enhancements to production of highly oxygenated molecules
  from autoxidation, Proc. Natl. Acad. Sci. U. S. A., 116(14), 6641–6646,
- 863 doi:10.1073/pnas.1810774116, 2019.
- 864 Pye, H. O. T., Nenes, A., Alexander, B., Ault, A. P., Barth, M. C., Simon, L., Collett, J. L.,
- 865 Fahey, K. M., Hennigan, C. J., Herrmann, H., Kanakidou, M., Kelly, J. T., Ku, I., Mcneill, V. F.,
- 866 Riemer, N., Shi, G., Tilgner, A., Walker, J. T., Wang, T., Weber, R., Xing, J., Zaveri, R. A. and
- 867 Zuend, A.: The Acidity of Atmospheric Particles and Clouds, Atmos. Chem. Phys, 20, 4809–
- 868 4888, doi:https://doi.org/10.5194/acp-20-4809-2020, 2020.
- 869 Randerson, J. T., van der Werf, G. R., Giglio, L., Collatz, G. J. and Kasibhatla, P. S.: Global Fire
- 870 Emissions Database, Version 4, (GFEDv4), ORNL DAAC, doi:10.3334/ORNLDAAC/1293,
  871 2015.
- 872 Ridley, D. A., Heald, C. L., Ridley, K. J. and Kroll, J. H.: Causes and consequences of
- 873 decreasing atmospheric organic aerosol in the United States, Proc. Natl. Acad. Sci. U. S. A.,
- 874 115(2), 290–295, doi:10.1073/pnas.1700387115, 2018.
- 875 Riedel, T. P., Lin, Y. H., Budisulistiorini, S. H., Gaston, C. J., Thornton, J. A., Zhang, Z.,
- 876 Vizuete, W., Gold, A. and Surratt, J. D.: Heterogeneous reactions of isoprene-derived epoxides:
- 877 Reaction probabilities and molar secondary organic aerosol yield estimates, Environ. Sci.
- 878 Technol. Lett., 2(2), 38–42, doi:10.1021/ez500406f, 2015.
- 879 Riedel, T. P., Lin, Y. H., Zhang, Z., Chu, K., Thornton, J. A., Vizuete, W., Gold, A. and Surratt,
- 880 J. D.: Constraining condensed-phase formation kinetics of secondary organic aerosol
- components from isoprene epoxydiols, Atmos. Chem. Phys., 16(3), 1245–1254,
- 882 doi:10.5194/acp-16-1245-2016, 2016.
- 883 Riva, M., Bell, D. M., Hansen, A. M. K., Drozd, G. T., Zhang, Z., Gold, A., Imre, D., Surratt, J.
- D., Glasius, M. and Zelenyuk, A.: Effect of Organic Coatings, Humidity and Aerosol Acidity on
   Multiphase Chemistry of Isoprene Epoxydiols, Environ. Sci. Technol., 50(11), 5580–5588,
- 886 doi:10.1021/acs.est.5b06050, 2016.
- 887 Riva, M., Chen, Y., Zhang, Y., Lei, Z., Olson, N. E., Boyer, H. C., Narayan, S., Yee, L. D.,
- 888 Green, H. S., Cui, T., Zhang, Z., Baumann, K., Fort, M., Edgerton, E., Budisulistiorini, S. H.,
- 889 Rose, C. A., Ribeiro, I. O., e Oliveira, R. L., dos Santos, E. O., Machado, C. M. D., Szopa, S.,
- 890 Zhao, Y., Alves, E. G., de Sá, S. S., Hu, W., Knipping, E. M., Shaw, S. L., Duvoisin Junior, S.,
- 891 de Souza, R. A. F., Palm, B. B., Jimenez, J.-L., Glasius, M., Goldstein, A. H., Pye, H. O. T.,
- 892 Gold, A., Turpin, B. J., Vizuete, W., Martin, S. T., Thornton, J. A., Dutcher, C. S., Ault, A. P.

- and Surratt, J. D.: Increasing Isoprene Epoxydiol-to-Inorganic Sulfate Aerosol Ratio Results in
- 894 Extensive Conversion of Inorganic Sulfate to Organosulfur Forms: Implications for Aerosol
- 895 Physicochemical Properties, Environ. Sci. Technol., 53(15), 8682-8694,
- 896 doi:10.1021/acs.est.9b01019, 2019.
- 897 Schmedding, R., Ma, M., Zhang, Y., Farrell, S., Pye, H. O. T., Chen, Y., Wang, C. tsan, Rasool,
- 898 Q. Z., Budisulistiorini, S. H., Ault, A. P., Surratt, J. D. and Vizuete, W.: A-Pinene-Derived
- 899 organic coatings on acidic sulfate aerosol impacts secondary organic aerosol formation from
- 900 isoprene in a box model, Atmos. Environ., 213(June), 456–462,
- 901 doi:10.1016/j.atmosenv.2019.06.005, 2019.
- 902 Schroder, J. C., Campuzano-Jost, P., Day, D. A., Shah, V., Larson, K., Sommers, J. M., Sullivan,
- 903 A. P., Campos, T., Reeves, J. M., Hills, A., Hornbrook, R. S., Blake, N. J., Scheuer, E., Guo, H.,
- 904 Fibiger, D. L., McDuffie, E. E., Hayes, P. L., Weber, R. J., Dibb, J. E., Apel, E. C., Jaeglé, L.,
- 905 Brown, S. S., Thornton, J. A. and Jimenez, J. L.: Sources and Secondary Production of Organic
- Aerosols in the Northeastern United States during WINTER, J. Geophys. Res. Atmos., 123(14),
   7771–7796, doi:10.1029/2018JD028475, 2018.
- 908 Schwantes, R. H., Charan, S. M., Bates, K. H., Huang, Y., Nguyen, T. B., Mai, H., Kong, W.,
- 909 Flagan, R. C. and Seinfeld, J. H.: Low-volatility compounds contribute significantly to isoprene
- 910 secondary organic aerosol (SOA) under high-NOx conditions, Atmos. Chem. Phys., 19(11),
- 911 7255–7278, doi:10.5194/acp-19-7255-2019, 2019.
- 912 Shah, V., Jaeglé, L., Jimenez, J. L., Schroder, J. C., Campuzano-Jost, P., Campos, T. L., Reeves,
- 913 J. M., Stell, M., Brown, S. S., Lee, B. H., Lopez-Hilfiker, F. D. and Thornton, J. A.: Widespread
- 914 Pollution From Secondary Sources of Organic Aerosols During Winter in the Northeastern
- 915 United States, Geophys. Res. Lett., 46(5), 2974–2983, doi:10.1029/2018GL081530, 2019.
  916 Shephard, M. W. and Cady-Pereira, K. E.: Cross-track Infrared Sounder (CrIS) satellite
- Shephard, M. W. and Cady-Pereira, K. E.: Cross-track Infrared Sounder (CrIS) satellite
   observations of tropospheric ammonia, Atmos. Meas. Tech., 8(3), 1323–1336, doi:10.5194/amt-
- 8-1323-2015, 2015.
   8-1323-2015, 2015.
- 919 Shrivastava, M., Cappa, C. D., Fan, J., Goldstein, A. H., Guenther, A. B., Jimenez, J. L., Kuang,
- 920 C., Laskin, A., Martin, S. T., Ng, N. L., Petaja, T., Pierce, J. R., Rasch, P. J., Roldin, P., Seinfeld,
- 921 J. H., Shilling, J., Smith, J. N., Thornton, J. A., Volkamer, R., Wang, J., Worsnop, D. R., Zaveri,
- 922 R. A., Zelenyuk, A. and Zhang, Q.: Recent advances in understanding secondary organic aerosol:
- 923 Implications for global climate forcing, Rev. Geophys., 55(2), 509–559,
- 924 doi:10.1002/2016RG000540, 2017.
- 925 Shrivastava, M., Andreae, M. O., Artaxo, P., Barbosa, H. M. J., Berg, L. K., Brito, J., Ching, J.,
- 926 Easter, R. C., Fan, J., Fast, J. D., Feng, Z., Fuentes, J. D., Glasius, M., Goldstein, A. H., Alves, E.
- 927 G., Gomes, H., Gu, D., Guenther, A., Jathar, S. H., Kim, S., Liu, Y., Lou, S., Martin, S. T.,
- 928 McNeill, V. F., Medeiros, A., de Sá, S. S., Shilling, J. E., Springston, S. R., Souza, R. A. F.,
- 929 Thornton, J. A., Isaacman-VanWertz, G., Yee, L. D., Ynoue, R., Zaveri, R. A., Zelenyuk, A. and
- 930 Zhao, C.: Urban pollution greatly enhances formation of natural aerosols over the Amazon
- 931 rainforest, Nat. Commun., 10(1), doi:10.1038/s41467-019-08909-4, 2019.
- 932 Silvern, R. F., Jacob, D. J., Kim, P. S., Marais, E. A., Turner, J. R., Campuzano-Jost, P. and
- 933 Jimenez, J. L.: Inconsistency of ammonium-sulfate aerosol ratios with thermodynamic models in
- 934 the eastern US: A possible role of organic aerosol, Atmos. Chem. Phys., 17(8), 5107–5118,
- 935 doi:10.5194/acp-17-5107-2017, 2017.
- 936 Solomon, P. A., Crumpler, D., Flanagan, J. B., Jayanty, R. K. M., Rickman, E. E. and McDade,
- 937 C. E.: U.S. National PM2.5 chemical speciation monitoring networks—CSN and IMPROVE:
- 938 Description of networks, J. Air Waste Manag. Assoc., 64(12), 1410–1438,

- 939 doi:10.1080/10962247.2014.956904, 2014.
- 940 Song, S., Gao, M., Xu, W., Shao, J., Shi, G., Wang, S., Wang, Y., Sun, Y. and McElroy, M. B.:
- 941 Fine-particle pH for Beijing winter haze as inferred from different thermodynamic equilibrium
- 942 models, Atmos. Chem. Phys., 18(10), 7423-7438, doi:10.5194/acp-18-7423-2018, 2018.
- 943 Surratt, J. D., Lewandowski, M., Offenberg, J. H., Jaoui, M., Kleindienst, T. E., Edney, E. O. and
- 944 Seinfeld, J. H.: Effect of acidity on secondary organic aerosol formation from isoprene, Environ.
- 945 Sci. Technol., (41), 5363-5369, doi:10.1021/es0704176, 2007a.
- 946 Surratt, J. D., Kroll, J. H., Kleindienst, T. E., Edney, E. O., Claeys, M., Sorooshian, A., Ng, N.
- 947 L., Offenberg, J. H., Lewandowski, M., Jaoui, M., Flagan, R. C. and Seinfeld, J. H.: Evidence for 948 organosulfates in secondary organic aerosol, Environ. Sci. Technol., (41), 517-527,
- 949 doi:10.1021/es062081q, 2007b.
- 950 Surratt, J. D., Chan, A. W. H., Eddingsaas, N. C., Chan, M., Loza, C. L., Kwan, A. J., Hersey, S.
- 951 P., Flagan, R. C., Wennberg, P. O. and Seinfeld, J. H.: Reactive intermediates revealed in
- 952 secondary organic aerosol formation from isoprene, Proc. Natl. Acad. Sci., 107(15), 6640-6645, 953 doi:10.1073/pnas.0911114107, 2010.
- 954 Tan, Y., Lim, Y. B., Altieri, K. E., Seitzinger, S. P. and Turpin, B. J.: Mechanisms leading to
- 955 oligomers and SOA through aqueous photooxidation: Insights from OH radical oxidation of
- 956 acetic acid and methylglyoxal, Atmos. Chem. Phys., 12(2), 801-813, doi:10.5194/acp-12-801-
- 957 2012, 2012.
- 958 Tao, Y. and Murphy, J. G.: The sensitivity of PM2.5 acidity to meteorological parameters and
- 959 chemical composition changes: 10-year records from six Canadian monitoring sites, Atmos. 960 Chem. Phys. Discuss., 1-21, doi:10.5194/acp-2019-238, 2019.
- 961 Travis, K. R., Jacob, D. J., Fisher, J. A., Kim, P. S., Marais, E. A., Zhu, L., Yu, K., Miller, C. C.,
- 962 Yantosca, R. M., Sulprizio, M. P., Thompson, A. M., Wennberg, P. O., Crounse, J. D., St Clair,
- 963 J. M., Cohen, R. C., Laughner, J. L., Dibb, J. E., Hall, S. R., Ullmann, K., Wolfe, G. M., Pollack,
- 964 I. B., Peischl, J., Neuman, J. A. and Zhou, X.: Why do models overestimate surface ozone in the
- 965 Southeast United States?, Atmos. Chem. Phys., 16(21), 13561-13577, doi:10.5194/acp-16-
- 966 13561-2016, 2016.
- 967 Tsui, W. G., Woo, J. L. and Mcneill, V. F.: Impact of Aerosol-Cloud Cycling on Aqueous
- 968 Secondary Organic Aerosol Formation, Atmosphere (Basel)., 10, 666,
- doi:doi:10.3390/atmos10110666, 2019. 969
- 970 Wagner, N. L., Brock, C. A., Angevine, W. M., Beyersdorf, A., Campuzano-Jost, P., Day, D., De
- 971 Gouw, J. A., Diskin, G. S., Gordon, T. D., Graus, M. G., Holloway, J. S., Huey, G., Jimenez, J.
- L., Lack, D. A., Liao, J., Liu, X., Markovic, M. Z., Middlebrook, A. M., Mikoviny, T., Peischl, 972
- 973 J., Perring, A. E., Richardson, M. S., Ryerson, T. B., Schwarz, J. P., Warneke, C., Welti, A.,
- 974 Wisthaler, A., Ziemba, L. D. and Murphy, D. M.: In situ vertical profiles of aerosol extinction,
- 975 mass, and composition over the southeast United States during SENEX and SEAC4RS:
- 976 Observations of a modest aerosol enhancement aloft, Atmos. Chem. Phys., 15(12), 7085–7102,
- 977 doi:10.5194/acp-15-7085-2015, 2015.
- 978 Weber, R. J., Guo, H., Russell, A. G. and Nenes, A.: High aerosol acidity despite declining
- 979 atmospheric sulfate concentrations over the past 15 years, Nat. Geosci., 9,
- 980 doi:10.1038/NGEO2665, 2016.
- 981 Woo, J. L. and McNeill, V. F.: simpleGAMMA v1.0 - A reduced model of secondary organic
- 982 aerosol formation in the aqueous aerosol phase (aaSOA), Geosci. Model Dev., 8(6), 1821–1829, doi:10.5194/gmd-8-1821-2015, 2015.
- 983
- 984 Worton, D. R., Surratt, J. D., Lafranchi, B. W., Chan, A. W. H., Zhao, Y., Weber, R. J., Park, J.

- 985 H., Gilman, J. B., De Gouw, J., Park, C., Schade, G., Beaver, M., Clair, J. M. S., Crounse, J.,
- 986 Wennberg, P., Wolfe, G. M., Harrold, S., Thornton, J. A., Farmer, D. K., Docherty, K. S.,
- 987 Cubison, M. J., Jimenez, J. L., Frossard, A. A., Russell, L. M., Kristensen, K., Glasius, M., Mao,
- 988 J., Ren, X., Brune, W., Browne, E. C., Pusede, S. E., Cohen, R. C., Seinfeld, J. H. and Goldstein,
- A. H.: Observational insights into aerosol formation from isoprene, Environ. Sci. Technol., (47),
   11403–11413, doi:10.1021/es4011064, 2013.
- 991 Xu, L., Suresh, S., Guo, H., Weber, R. J. and Ng, N. L.: Aerosol characterization over the
- 992 southeastern United States using high-resolution aerosol mass spectrometry: Spatial and seasonal
- 993 variation of aerosol composition and sources with a focus on organic nitrates, Atmos. Chem.
- 994 Phys., 15(13), 7307–7336, doi:10.5194/acp-15-7307-2015, 2015a.
- 995 Xu, L., Guo, H., Boyd, C. M., Klein, M., Bougiatioti, A., Cerully, K. M., Hite, J. R., Isaacman-
- 996 VanWertz, G., Kreisberg, N. M., Knote, C., Olson, K., Koss, A., Goldstein, A. H., Hering, S. V.,
- 997 de Gouw, J., Baumann, K., Lee, S.-H., Nenes, A., Weber, R. J. and Ng, N. L.: Effects of
- 998 anthropogenic emissions on aerosol formation from isoprene and monoterpenes in the
- southeastern United States, Proc. Natl. Acad. Sci., 112(1), 37–42, doi:10.1073/pnas.1417609112,
  2015b.
- 1001 Xu, L., Middlebrook, A. M., Liao, J., de Gouw, J. A., Guo, H., Weber, R. J., Nenes, A., Lopez-
- 1002 Hilfiker, F. D., Lee, B. H., Thornton, J. A., Brock, C. A., Neuman, J. A., Nowak, J. B., Pollack,
- 1003 I. B., Welti, A., Graus, M., Warneke, C. and Ng, N. L.: Enhanced formation of isoprene-derived 1004 organic aerosol in sulfur-rich power plant plumes during Southeast Nexus, J. Geophys. Res.,
- 1005 121, 11137–11153, doi:10.1002/2016JD025156, 2016.
- 1006 Xu, L., Pye, H. O. T., He, J., Chen, Y., Murphy, B. N. and Ng, N. L.: Experimental and model
- estimates of the contributions from biogenic monoterpenes and sesquiterpenes to secondary
  organic aerosol in the southeastern United States, Atmos. Chem. Phys., 18(17), 12613–12637,
- 1009 doi:10.5194/acp-18-12613-2018, 2018.
- 1010 Yu, K., Jacob, D. J., Fisher, J. A., Kim, P. S., Marais, E. A., Miller, C. C., Travis, K. R., Zhu, L.,
- 1011 Yantosca, R. M., Sulprizio, M. P., Cohen, R. C., Dibb, J. E., Fried, A., Mikoviny, T., Ryerson, T.
- 1012 B., Wennberg, P. O. and Wisthaler, A.: Sensitivity to grid resolution in the ability of a chemical
- transport model to simulate observed oxidant chemistry under high-isoprene conditions, Atmos.
  Chem. Phys., 16(7), 4369–4378, doi:10.5194/acp-16-4369-2016, 2016.
- 1015 Zhang, H., Yee, L. D., Lee, B. H., Curtis, M. P., Worton, D. R., Isaacman-VanWertz, G.,
- 1016 Offenberg, J. H., Lewandowski, M., Kleindienst, T. E., Beaver, M. R., Holder, A. L., Lonneman,
- 1017 W. A., Docherty, K. S., Jaoui, M., Pye, H. O. T., Hu, W., Day, D. A., Campuzano-Jost, P.,
- 1018 Jimenez, J. L., Guo, H., Weber, R. J., de Gouw, J., Koss, A. R., Edgerton, E. S., Brune, W.,
- 1019 Mohr, C., Lopez-Hilfiker, F. D., Lutz, A., Kreisberg, N. M., Spielman, S. R., Hering, S. V.,
- 1020 Wilson, K. R., Thornton, J. A. and Goldstein, A. H.: Monoterpenes are the largest source of
- 1021 summertime organic aerosol in the southeastern United States, Proc. Natl. Acad. Sci.,
- 1022 201717513, doi:10.1073/pnas.1717513115, 2018a.
- 1023 Zhang, Y., Chen, Y., Lambe, A. T., Olson, N. E., Lei, Z., Craig, R. L., Zhang, Z., Gold, A.,
- 1024 Onasch, T. B., Jayne, J. T., Worsnop, D. R., Gaston, C. J., Thornton, J. A., Vizuete, W., Ault, A.
- 1025 P. and Surratt, J. D.: Effect of the Aerosol-Phase State on Secondary Organic Aerosol Formation
- 1026 from the Reactive Uptake of Isoprene-Derived Epoxydiols (IEPOX), Environ. Sci. Technol.
- 1027 Lett., 5(3), 167–174, doi:10.1021/acs.estlett.8b00044, 2018b.
- 1028 Zhang, Y., Chen, Y., Lei, Z., Olson, N. E., Riva, M., Koss, A. R., Zhang, Z., Gold, A., Jayne, J.
- 1029 T., Worsnop, D. R., Onasch, T. B., Kroll, J. H., Turpin, B. J., Ault, A. P. and Surratt, J. D.: Joint
- 1030 Impacts of Acidity and Viscosity on the Formation of Secondary Organic Aerosol from Isoprene

- 1031 Epoxydiols (IEPOX) in Phase Separated Particles, ACS Earth Sp. Chem., 3(12), 2646–2658,
- 1032 doi:10.1021/acsearthspacechem.9b00209, 2019a.
- 1033 Zhang, Y., Nichman, L., Spencer, P., Jung, J. I., Lee, A., Heffernan, B. K., Gold, A., Zhang, Z.,
- 1034 Chen, Y., Canagaratna, M. R., Jayne, J. T., Worsnop, D. R., Onasch, T. B., Surratt, J. D.,
- 1035 Chandler, D., Davidovits, P. and Kolb, C. E.: The Cooling Rate- And Volatility-Dependent
- 1036 Glass-Forming Properties of Organic Aerosols Measured by Broadband Dielectric Spectroscopy,
- 1037 Environ. Sci. Technol., 53(21), 12366–12378, doi:10.1021/acs.est.9b03317, 2019b.
- 1038 Zheng, Y., Unger, N., Hodzic, A., Emmons, L., Knote, C., Tilmes, S., Lamarque, J. F. and Yu,
- 1039 P.: Limited effect of anthropogenic nitrogen oxides on secondary organic aerosol formation,
- 1040 Atmos. Chem. Phys., 15(23), 13487–13506, doi:10.5194/acp-15-13487-2015, 2015.
- 1041 Ziemann, P. J. and Atkinson, R.: Kinetics, products, and mechanisms of secondary organic
- 1042 aerosol formation, Chem. Soc. Rev., 41(19), 6582, doi:10.1039/c2cs35122f, 2012.
- 1043

1044

1045



**Figure 1.** Comparison of June-July-August averaged surface OA concentration ( $\mu g/m^3$ ) over the southeast US between the default model and the observation from IMPROVE and SEARCH network. Colored shades represent different components of modeled OA. IEPOX-, glyoxal-, and other isoprene-SOA are from aqueous uptake of isoprene oxidation products. Terpene- and anthropogenic SOA are dry SOA calculated using volatility-basis-set.



**Figure 2. (A)** Monthly surface OA and sulfate  $(SO_4^{2-})$  concentration  $(\mu g/m^3)$  averaged over the southeast US from IMPROVE, SEARCH and the default model. **(B)** Monthly surface concentrations of IEPOX-SOA and the sum of POA and dry SOA from the default model, and IEPOX-SOA from the CT H01 simulation.



**Figure 3.** Standardized monthly surface IEPOX-SOA concentration, sulfate concentration, aerosol H<sup>+</sup> activity and isoprene emission from the default model. All variables are averaged over the southeast US, and have been divided by 1 standard deviations, therefore are unitless.



**Figure 4**. (A) Schematic diagram of IEPOX reactive uptake coefficient ( $\gamma_{IEPOX}$ ). Colored lines indicate the position of JJA-averaged organic mass fraction and aerosol H<sup>+</sup> activity in 2000-2013 from the 'Default', 'CT' and 'CT\_H01' simulations. (B) Simulated aerosol acidity (mol/L) from the default, 'CT\_newNH<sub>3</sub>' and 'CT\_H01' simulations. (C) JJA-averaged surface OA ( $\mu$ g/m<sup>3</sup>) from IMPROVE, SEARCH and all model simulations. (D) Standard deviation of OA ( $\mu$ g/m<sup>3</sup>) between June, July and August from IMPROVE, SEARCH and all model simulations. All results are averaged over the southeast US.



**Figure 5.** Relationships between monthly OA and sulfate concentrations ( $\mu g/m^3$ ). Each dot represents monthly data averaged from all sites from each network within the southeast US.

# **Supplementary Information**

**Table S1.** Multivariate linear regression analysis of modeled IEPOX-SOA.  $\beta_{1,}\beta_{2}$  and  $\beta_{3}$  are standardized partial regression coefficients associated with sulfate, aerosol H<sup>+</sup> activity and isoprene emission. R<sup>2</sup> represent the square of correlation between IEPOX-SOA and the linear-fitted function. The  $r_{1}^{2}$ ,  $r_{2}^{2}$  and  $r_{3}^{2}$  within parentheses are square of direct correlation between IEPOX-SOA and each variable.

	Sulfate $\beta_1 (r_1^2)$	$H^+$ activity $oldsymbol{\beta}_2$ ( $r_2^2$ )	ISOP emission $\beta_3 (r_3^2)$	R <sup>2</sup>
2000-2013	0.39 (0.64)	0.50 (0.71)	0.34 (0.18)	0.88
2000-2004	0.46 (0.76)	0.43 (0.64)	0.25 (0.55)	0.93
2005-2008	0.34 (0.62)	0.57 (0.82)	0.31 (0.24)	0.94
2009-2013	0.45 (0.55)	0.27 (0.56)	0.42 (0.53)	0.84



**Figure S1.** OA concentration ( $\mu$ g/m<sup>3</sup>) averaged in JJA 2000-2013 from the default model and the IMPROVE and SEARCH networks. Rectangle defines the SEUS region. Circles and triangles represent sites from IMPROVE and SEARCH, respectively. Black dots represent the location of CSN sites, but due to the discontinuity of CSN organic carbon measurements, the 2000-2013 OA concentrations are not shown for CSN sites.



Figure S2. Monthly observed OA from the CSN networks. Units are  $\mu g/m^3$ .



**Figure S3.** Monthly NH<sub>3</sub> emissions from the default NEI11v1 emission inventory and the CrISderived emission inventory used in the 'CT\_NH<sub>3</sub>' sensitivity simulation. Emissions are averaged in 2000-2013.



**Figure S4.** Monthly surface aerosol H<sup>+</sup> activity (mol/L), NH<sub>3</sub> emission, SO<sub>2</sub> emission (mg/m<sup>2</sup>/hr) and aerosol  $NH_4^+/SO_4^{2-}$  ratio from the default and 'CT\_newNH<sub>3</sub>' simulations.



Figure S5. The monthly surface IEPOX-SOA concentration ( $\mu g/m^3$ ) over the SEUS from all simulations.



**Figure S6.** Comparison of **(A)** IEPOX-SOA, **(B)** OA between AMS measurement and the default and CT\_H01 simulations in the SOAS campaign from June 01 to July 15, 2013. In (A), IEPOX-SOA is obtained by a Positive Matrix Factorization (PMF) analysis of the AMS-measured OA. Units are  $\mu g/m^3$ .