



| 1 | The Effect of Meteorological Conditions and Atmospheric |
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| 2 | Composition in the Occurrence and Development of New Particle |
| 3 | Formation (NPF) Events in Europe |
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ABSTRACT

Although new particle formation (NPF) events have been studied extensively for some decades, the 39 mechanisms that drive their occurrence and development are yet to be fully elucidated. Laboratory 40 41 studies have done much to elucidate the molecular processes involved in nucleation, but this knowledge has yet to be linked to NPF events in the atmosphere, except at very clean air sites. 42 43 There is great difficulty in successful application of the results from laboratory studies to real atmospheric conditions, due to the diversity of atmospheric conditions and observations found, as 44 45 NPF events occur almost everywhere in the world without following a clearly defined trend of 46 frequency, seasonality, atmospheric conditions or event development. The present study seeks common features in nucleation events by applying a binned linear regression over an extensive 47 dataset from 16 sites of various types (rural and urban backgrounds as well as roadsides) in Europe. 48 49 A clear positive relation is found between the solar radiation intensity, temperature and atmospheric pressure with the frequency of NPF events, while relative humidity presents a negative relation with 50 NPF event frequency. Wind speed presents a less consistent relationship which appears to be 51 52 heavily affected by local conditions. While some meteorological variables appear to have a crucial effect on the occurrence and characteristics of NPF events, especially at rural sites, it appears that 53 their role becomes less marked when at higher values. 54 55 The analysis of chemical composition data presents interesting results. Concentrations of almost all 56 57 chemical compounds studied (apart from O₃) and the Condensation Sink (CS) have a negative





relation with NPF event probability, though areas with higher average concentrations of SO₂ had higher NPF event probability. Particulate Organic Carbon (OC), Volatile Organic Compounds (VOCs) and particulate phase sulphate consistently had a positive relation with the growth rate of the newly formed particles. As with some meteorological variables, it appears that at increased concentrations of pollutants or the CS, their influence upon NPF probability is reduced.

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1. INTRODUCTION

New Particle Formation (NPF) events are an important source of particles in the atmosphere 65 (Merikanto et al., 2009; Spracklen et al., 2010), which are known to have adverse effects on human 66 67 health (Schwartz et al., 1996; Politis et al., 2008; Kim, et al., 2015) as well as affecting the optical and physical properties of the atmosphere (Makkonen et al., 2012; Seinfeld and Pandis, 2012). 68 69 While they occur almost everywhere in the world (Dall'Osto et al., 2018; Kulmala et al., 2017; 70 O'Dowd et al., 2002; Wiedensohler et al., 2019; Chu et al., 2019; Kerminen et al., 2018), great 71 diversity is found in the atmospheric conditions within which they take place. Many studies have 72 been done in a large number of different types of locations (urban, traffic, regional background) around the world and differences were found in both the seasonality and intensity of NPF events. 73 74 To an extent this variability is due to the mix of conditions that are specific to each location, which 75 blurs the general understanding of the conditions that are favourable for the occurrence of NPF events (Berland et al., 2017; Bousiotis et al., 2020). For example, solar radiation is considered as 76 one of the most important factors in the occurrence of NPF events (Kulmala and Kerminen, 2008; 77 78 Kürten et al., 2016; Pikridas et al., 2015; Salma et al., 2011), as it is needed for the photochemical reactions that lead to the formation of sulphuric acid (Petäjä et al., 2009; Cheung et al., 2013), 79 80 which is considered as the main component of the formation and growth of the initial clusters (Iida 81 et al., 2008; Weber et al., 1995); although in many cases, NPF events did not occur in the seasons 82 with the highest insolation (Park et al., 2015; Vratolis et al., 2019). Similarly, higher temperatures are considered favourable for the growth of the newly formed particles as increased concentrations 83





84 of both Biogenic Volatile Organic Compounds (BVOCs) and Anthropogenic Volatile Organic 85 Compounds (AVOCs) (Yamada, 2013; Paasonen et al., 2013) and their oxidation products (Ehn et al., 2014) are associated with the growth of the particles. This appears to be true in most cases, as 86 87 higher growth rates are found in most cases in the local summer (Nieminen et al., 2018), although the actual importance of those VOCs in the occurrence of NPF events is still not fully elucidated. 88 89 The effect of other meteorological variables is even more complex, with studies presenting mixed results on the effect of the wind speed and atmospheric pressure. Extreme values of those variables 90 91 may be favourable for the occurrence of NPF events, as they are associated with increased mixing 92 in the atmosphere, but at the same time suppress due to increased dilution of precursors (Brines et al., 2015; Rimnácová et al., 2011; Shen et al., 2018; Siakavaras et al., 2016), or favour them due to 93 94 a reduced condensation sink (CS). 95 The effect of atmospheric composition on NPF events is also a puzzle of mixed results. While the 96 negative effect of the increased CS is widely accepted (Kalkavouras et al., 2017; Kerminen et al., 97 98 2004; Wehner et al., 2007), cases are found when NPF events occur on days with higher CS 99 compared to average conditions (Größ et al., 2018; Kulmala et al., 2005). Sulphur dioxide (SO₂), 100 which is one of the most important contributors to many NPF pathways, in most studies was found 101 in lower concentrations on NPF event days compared to average conditions (Alam et al., 2003; 102 Bousiotis et al., 2019), although there are studies that have reported the opposite (Woo et al., 2001; Charron et al., 2008). Additionally, in a combined study of NPF events in China, events were found 103





104 to be more probable under sulphur-rich conditions rather than sulphur-poor (Jayaratne et al., 2017). Similar is the case with the BVOCs and AVOCs, which present great variability depending the area 105 studied (Dai et al., 2017), and their contribution in the growth of the particles is not fully understood 106 107 yet. Until recently, it was considered unlikely for NPF events as they are considered in the present study, deriving from secondary formation not associated with traffic related processes such as 108 109 dilution of the exhaust, to occur within the complex urban environment due to the increased 110 presence of compounds, mainly associated with combustion processes, which would suppress the 111 survival of the newly formed particles within this type of environment (Kulmala et al., 2017). 112 Despite that though, NPF events were found to occur within even the most polluted areas and sometimes with high formation and growth rates (Bousiotis et al., 2019; Yao et al., 2018). 113 114 It is evident that while a general knowledge of the role of the meteorological and atmospheric variables has been achieved, there is great uncertainty over the extent and variability of their effect 115 116 (and for some of them even their actual effect) in the mechanisms of NPF in real atmospheric 117 conditions, especially in the more complex urban environment (Harrison, 2017). The present study, 118 using an extensive dataset from 16 sites in six European countries, attempts to elucidate the effect 119 of several meteorological and atmospheric variables not only in general, but also depending on the 120 geographical region or type of environment. While studies with multiple sites have been reported in 121 the past, to our knowledge this is the first study that focuses directly on the effect of these variables 122 upon the probability of NPF events as well as the formation and growth rates of newly formed particles in real atmospheric conditions. 123





124 2. DATA AND METHODS

125 2.1 Site Description and Data Availability

The present study uses a total of more than 85 years of hourly data from 16 sites from six countries of Europe of various land usage and climates from which 1950 NPF events were extracted and studied. A list of the available data and a brief description for each site is found in Table 1 (for the ease of reading the sites are named by the country of the site followed by the last two letters which refer to the type of site, being RU for rural/regional background, UB for urban background and RO for roadside), while a map of the sites is found in Figure 1. The NPF frequency and formation rate for each site is found in Table 2.

2.2 Methods

135 2.2.1 NPF events selection

NPF events were selected using the method proposed by Dal Maso et al (2005). As of this, an NPF event is considered when a new mode of particles appears in the nucleation mode, prevails for some hours and shows signs of growth. The events can then be classified into classes I and II according to the level of confidence, while class I events can be further classified to Ia and Ib, with Ia events having both a clear formation of a by new mode of particles as well as a distinct growth of the new mode of particles, while Ib consists of rather clear events that fail though by at least one of the criteria set. In the present study, only the events of class Ia were considered with the additional criterion of at least 1 nm h⁻¹ growth for at least 3 hours.





144 2.2.2 Calculation of condensation sink, growth rate, formation rate, and NPF event

- 145 **probability**
- 146 The condensation sink (CS) is calculated according to the method proposed by Kulmala et al.,
- 147 (2001) as:

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$$CS = 4\pi D_{vap} \sum \beta_M r N$$

- 151 where r and N is the radius and number concentration of the particles and D_{vap} is the diffusion
- 152 coefficient calculated as (Poling et al., 2001):

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$$D_{vap} = 0.00143 \cdot T^{1.75} \frac{\sqrt{M_{air}^{-1} + M_{vap}^{-1}}}{P\left(D_{x,air}^{\frac{1}{3}} + D_{x,vap}^{\frac{1}{3}}\right)^2}$$

- for T = 293 K and P = 1013.25 mbar. M and D_x are the molar mass and diffusion volume for air and
- sulphuric acid. β_M is the Fuchs correction factor calculated as (Fuchs and Sutugin, 1971):

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$$\beta_{M} = \frac{1 + K_{n}}{1 + \left(\frac{4}{3a} + 0.377\right)K_{n} + \frac{4}{3a}K_{n}^{2}}$$

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where K_n is the Knudsen number, calculated as $K_n = 2\lambda_m/d_p$ where λ_m is the mean free path of the

162 gas.

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164 Growth rate (GR) is calculated as (Kulmala et al., 2012):

166 GR = $\frac{D_{P_2} - D_{P_1}}{t_2 - t_1}$

168 for the size range between the minimum available particle diameter up to 30 nm (50 nm for the UK

sites due to the higher minimum particle size available). The time window used for the calculation

170 of the growth rate was from the start of the event until a) growth stopped, b) GMD reached the

171 upper limit set or c) the day ended.

173 The formation rate J was calculated using the method proposed by (Kulmala et al., 2012) as:

175 $J_{d_p} = \frac{dN_{d_p}}{dt} + CoagS_{d_p} \times N_{d_p} + \frac{GR}{\Delta d_p} \times N_{d_p} + S_{losses}$

where CoagS_{dp} is the coagulation rate of particles of diameter d_p, calculated as (Kerminen et al.,

178 2001):

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$$180 \quad CoagS_{d_p} = \int K(d_p, d'_p) \, n(d'_p) dd'_p \, \cong \, \sum_{d'_p = d_p}^{d'_p = max} K(d_p, d'_p) \, N_{d_p}$$

182 $K(d_p, d'_p)$ is the coagulation coefficient of particles with diameters d_p and d'_p , while S_{losses} accounts

183 for additional loss terms (i.e. chamber wall losses), which are not applicable in the present study.

For the present study, the formation rate of particles of diameter of 10 nm was calculated for

uniformity (16 nm for the UK sites), though most sites had data for particle sizes below 10 nm.

187 The NPF probability was calculated by the number of NPF event days divided by the number of

days with available data in the given group (temporal, wind direction sector etc.). The results

189 presented in this study were also normalised according to the data availability, as:

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$$NPF_{probability} = \frac{N_{NPF \ event \ days \ for \ group \ of \ days \ X}}{N_{days \ with \ available \ data \ for \ group \ of \ days \ X}}$$

2.2.3 Calculation of the slope and intercept for the variables used

Due to the large datasets available and the great spread of the values, a direct comparison between a given variable and any of the characteristics associated with NPF events (NPF probability, growth rate and formation rate) always provided results with low significance. As a result, an alternative method which can provide a reliable result without the noise of the large datasets was used in the

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present study, to investigate the relations between the variables which are considered to be associated with the NPF events. For this, a timeframe which is more directly associated with the NPF events typically observed in the mid-latitudes was chosen. For NPF probability and GR the timeframe between 05:00 to 17:00 LT was chosen, which is considered the time when the vast majority of NPF events take place and further develop with the growth of the particles. For the formation rate a smaller timeframe was chosen, 09:00 to 15:00 LT (Local Time) which is \pm 3 hours from the time of the maximum formation rate found for almost all sites (12:00 LT). This was done to exclude as far as possible the effect of the morning rush at the roadsides, as well as only to include the time window when the formation rate is mostly relevant to NPF events (negative values that are more probable outside this timeframe would bias the results). Specifically, for the CS the timeframe 05:00 to 10:00 LT was chosen. This was done to avoid including the direct effect of the NPF events as well as to provide results for the conditions which either promote or suppress the characteristics studied, which specifically for the CS are more important before the start of the events. The extreme values (very high or very low) which bias the results only carrying a very small piece (forming bins of very small size) of information were then removed, though 90% of the available data were used for all the variables. The data left was separated into smaller bins and a minimum of 10 bins was required for each variable (for example if the difference between the minimum and the maximum relative humidity (RH) is 70%, then 14 bins





each with a range of 5% were formed). The variables of interest were then averaged for each binand plotted, and a linear relation was considered for each one of them.

The slope of the linear relations (a_{N,a_G} and a_{J} for NPF probability, growth rate and formation rate J_{10} accordingly) found in this analysis should be used with great caution as apart from the atmospheric conditions (local and meteorological as well as atmospheric composition) it is also affected by the variable in question (e.g. a greater NPF probability will provide a greater slope), resulting in giving the same trend for all the atmospheric variables tested; the sites with the higher values of these variables (NPF probability and formation rate) always had greater slope values and vice versa. In order to remove the effect of the variable in question (NPF probability or formation rate – growth rate will provide an untrustworthy result as it is calculated in a different range for each site due to the lower available size of particles), the slopes were normalised by dividing them by their respective variable (e.g. divide the slope of the NPF probability with the NPF probability), providing with a new normalised slope (a_{N} * for NPF probability or a_{J} * for the formation rate) that will have no significance other than its absolute value, which can be used for direct comparisons:

$$a_{N}^{*} = \frac{a_{N}}{NPF \%}$$

Where a_N is the slope of the relation between the given variable and NPF probability (NPF %)

$$a_J^* = \frac{a_J}{J_{10}}$$





236 Where a_J is the slope of the relation between the given variable and the formation rate of 10 nm particles J_{10} (J_{16} for the UK sites). 237 238 239 **3.** RESULTS In this study NPF events are generally observed as particles grow from a smaller size (typically 3-240 15 nm depending on the size detection limit of instruments used) to 30 nm or larger. They therefore 241 242 reflect the result both of nucleation, which creates new particles of 1-2 nm (not detected with the instruments used in this study), and growth to larger sizes. In analysing NPF events, we therefore 243 244 consider three diagnostic features: the frequency of events occurring (i.e. days with an event divided by total days with relevant 245 246 data), the rate of particle formation at a given size (J_{10} in this case), 247 248 the growth rate of particles from the lower measurement limit to 30 nm (or 50 nm for the UK 249 sites). 250 251 3.1 **Meteorological Conditions** The slopes and R² from the analysis of the meteorological variables, as well as the average 252 conditions of these variables are found in Table 3. The results for each site and variable are found in 253 254 Figure S1.



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3.1.1 Solar radiation intensity

As mentioned earlier, solar radiation is considered as one of the most important variables in NPF occurrence, as it contributes to the production of H₂SO₄ which is a main component of the initial clusters and participates in the early growth of the newly formed particles. Hidy et al. (1994) reported up to six times higher SO₂ oxidation rates into H₂SO₄ in typical summer conditions compared to winter). For almost all sites this relation is confirmed with very strong correlations between the intensity of solar radiation and the probability for NPF events. The relation between the solar radiation and NPF probability was positive at all sites and only three sites (FINUB, SPARU and GREUB) presented weak correlations (R² below 0.40). Weaker correlations were found for the southern European sites, which might be associated with the higher averages for solar radiation, or the interference of other processes (such as coinciding with increased CS by recirculation of air masses (Carnerero et al., 2019), possibly making it less of an important factor for these areas. The relationship of solar radiation to the growth rate was weaker in all cases and did not present a clear trend. A few sites presented a strong correlation, which in all cases were background sites (either rural or urban). The relation found in most cases was positive apart from two roadsides and GREUB, though due to the low R² these results cannot be used with confidence. It seems though that the solar radiation intensity is probably a more important factor at background sites rather than at roadsides, where possibly local conditions (such as local emissions) are more important. Finally, the formation rate has a positive relation with the solar radiation intensity, with strong correlations





in most areas. The correlations were stronger at the rural background sites compared to the roadsides, which further underlines the increased importance of this factor at this type of site. A negative relation between the solar radiation intensity and the formation rate was found at the GRERU site but the R² is very low.

Plotting the normalised slopes for NPF event probability a_N^* with the average solar radiation at each site (Figure 2) a negative relation is found ($R^2 = 0.62$), with the southern areas (those with higher average solar intensity) having smaller a_N^* compared to those in higher latitudes (and thus with a lower average solar radiation). This may indicate that while solar radiation is a deciding factor in the occurrence of an NPF event, when in greater intensity its role becomes relatively less important, a finding that was also implied by Wonaschütz et al. (2015). Additionally, the a_J^* was found to be higher at all rural sites compared to their respective roadsides (and urban background sites for all but the Greek and German ones), making it a more important factor at this type of site (Figure 3).

3.1.2 Relative humidity

Relative humidity is considered to have a negative effect on the occurrence of NPF events (Jeong et al., 2010; Hamed et al., 2011; Park et al., 2015; Dada et al., 2017; Li et al., 2019). While water in the atmosphere is one of the main compounds needed for the formation of the initial clusters either on the binary or ternary nucleation theory (Korhonen et al., 1999; Mirabel and Katz, 1974), in atmospheric conditions it may also play a negative role suppressing the number concentrations of

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new particles by increasing aerosol surface area. Consistent with this, a negative relation of the RH with NPF probability was found for all the sites of this study with very high R² for almost all of them. This is not simple to interpret as solar radiation, temperature, RH and CS are not independent variables, since an increase in temperature of an air mass due to increased solar radiation will be associated with reduced RH, which in turn affects the CS. The sites in Greece presented lower R² compared to the other sites while, GRERU was found to have the weakest correlation. Growth rate on the other hand had a variable relation, either positive or negative, with only a handful of background sites having strong correlations. Among these the German background sites as well as FINRU, which were among the sites with the highest average RH (average RH for GERRU is 81.9%, GERUB is 78.7% and FINUB is 80.1%) presented a negative relation between the RH and growth rate, while DENRU (average RH at 75.7%) had a positive relation, which might indicate that the relation between these two variables may vary depending upon the RH range. Formation rate also appears to have a negative relation with the RH, though this relation was significant (R^2) 0.40) for only 6 sites, which once again in most cases are sites with higher RH average conditions. Along with the results of the growth rate this might indicate that the RH becomes a more important factor in the development of NPF events as its values increase.

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The normalised slopes once again provide some additional information. Regarding the NPF probability, it is found that the a_N^* was more negative at rural sites compared to roadsides. This indicates that the RH has a smaller effect at roadsides, as other variables, such as the atmospheric





316 composition, are probably more important within the complex environment in this type of sites.

317 Additionally, the relation between a_N* and average RH at the sites had a negative relation (R² = 0.46), which further shows that the RH becomes a more important factor at higher values (Figure 4). Furthermore, at the sets of rural and roadside sites with R² higher than 0.40 for the relation 320 between RH and the formation rate (UK and German sites), it was found that the a_J* was more 321 negative at the rural sites which indicates that the RH is a more important factor at rural sites 322 compared to their respective roadsides.

3.1.3 Temperature

Temperature can have both a direct and indirect effect in the development of NPF events, as it is directly associated with the abundance of biogenic volatile carbon which is an important group of compounds whose oxidation products can participate in nucleation itself (Lehtipalo et al., 2018; Rose et al., 2018), as well as in the growth of newly formed particles, while it may affect the particle size distributions or number concentrations through other processes such as particle evaporation. Most of the sites of the present study presented a strong relation of NPF probability with temperature, which in most cases was positive, though in many cases (such as the Danish, Spanish and Finnish sites) there seems to be a peak in the NPF probability at some temperature, after which a decline starts (though being at the higher end does not greatly affect the results). Sites with smaller R² (weaker association with temperature), were mainly those that have a seasonal variation that favoured seasons other than summer. These sites not only had weaker relation of NPF



336 probability with temperature, but in most cases had a negative relation (background sites in Finland, 337 Spain and Greece). The Finnish sites, having the lowest average temperatures and a sufficient amount of data below zero temperature, show at all three sites the possible presence of a peak in the 338 339 NPF event probability for temperatures below zero. This seems to be the cause of the weak relations found there and they seem to be associated with the formation rate J₁₀, which also seems to have an 340 341 increasing trend below zero degrees. This may be the result of increased stability of molecular 342 clusters at lower temperatures, as well as the possible enhancement of growth mechanisms in lower 343 temperatures (below 5°C) by other chemical compounds in the atmosphere (i.e. nitric acid and ammonia) as found by Wang et al., (2020). Laboratory experiments show that the characteristics of 344 345 organic aerosol forming from alpha-pinene is governed by gas phase oxidation (e.g. Ye et al. 2019). In the real atmosphere, the higher temperature enhances the amount of biogenic vapours (e.g. 346 Paasonen et al. 2013), and although the oxidation can be more efficient in higher temperatures, the 347 348 lower temperatures favour formation of more non-volatile compounds (Ye et al. 2019; Stolzenburg 349 et al. 2018). 350 Growth rate had a more uniform trend, with almost all sites having a positive relation with 351 temperature (apart from GERRO, though with $R^2 = 0.00$). This relation was very strong for most 352 353 sites, which is also confirming the summer peak found for the growth rate at most of these sites. A strong relation with temperature was also found for the formation rate for most sites, and was 354

positive for almost all sites (apart from FINRO with $R^2 = 0.01$ and the Greek sites). As with the





NPF probability, in general the sites with a seasonal variation of events that favoured summer had the strongest relation (high R^2) of the formation rate with temperature, which might indicate that this variable, either through its direct or indirect effect is an important one for the seasonal variability of NPF events in a given area.

The normalised slopes for this variable did not present a clear trend among the areas studied, other than presenting greater a_N^* for the sites with a summer peak in their NPF event seasonal variation. As with other meteorological variables, the importance of this variable became smaller with increased values in the average conditions for both the NPF probability (Figure 5) and J_{10} , though these relations were not significant (biased by the very low average temperatures and different behaviour of the variables at the Finnish sites, without which the relation becomes a lot clearer as pointed in Figure S2). The variation though within the sites of the same area (different sites in same country / region) appears to directly follow the variability of temperature, showing that the temperature directly affects the occurrence of NPF events when other factors remain constant, having a negative trend for all countries but Finland. The a_1^* though is found to be greater (positively or negatively) at the rural background sites than at the other two types of sites at all areas studied, showing that it is a more important factor for the formation rate at this type of site compared to others (Figure 6).



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3.1.4 Wind speed

Wind speed may have both a positive and a negative effect on the occurrence of NPF events. On one hand, it may promote NPF events by the increased mixing of the condensable compounds in the atmosphere as well as by reducing the CS, while on the other hand high wind speeds may suppress NPF events due to increased dilution. It should be considered that the variability found is also affected by the specific conditions found at each site. The wind speed measurements in many cases, especially in urban sites, can be biased by the local topography or specific conditions found at each site, thus representing the local conditions for this variable rather than the regional ones. Similarly, measurements of wind speed at well sited meteorological stations may be more representative of regional conditions, than of those affecting the sites of nucleation measurement. The sites in this study presented mixed results, both in the importance as well as the effect of the wind speed variability. Three different behaviours were found in the variation of NPF event probability and wind speed which appear to be associated with local conditions as they are almost uniformly found among the sites within close proximity. Some sites presented a steady increase of NPF event probability with wind speed (Danish sites as well as UKUB, FINRU, SPAUB and GRERU), while others were found to steadily decline with increasing wind speeds (German sites – it should be noted that the German sites are the only ones that are located at a great distance from the sea), while some were found to reach a peak and then decline, which also leads to smaller R² (UKRU, UKRO, SPARU and to a lesser extent GREUB). The reasons for these differences between the sites are very hard to distinguish as apart from the wind speed the origin and the characteristics of these air





396 masses play a crucial role. Following this, it appears that NPF probability is very low or zero for 397 wind speeds close to calm for the sites with an increasing trend (as well as those that have a peak and decline after), while the opposite is observed for the German sites where the maximum NPF 398 399 probability is found for very low wind speeds. 400 Similarly, the effect of different wind speeds upon the growth rate also varied a lot, though it was 401 found to be negative in all the cases where R² was higher than 0.50 (UKUB, DENRU, DENRO, 402 403 GERRU, GERUB and GREUB). Finally, the formation rate was found to have a significant 404 correlation only at two sites (UKRO and DENRU), probably indicating that the variability of the 405 wind speed either does not affect this variable or its effect is rather small. 406 407 The normalised slopes did not have any notable relation to either the NPF probability or the 408 formation rate further confirming that the effect of the different wind speeds is not due to its 409 variability only, but it is also influenced by the characteristics of the incoming air masses as well as 410 specific local conditions found at each site. 411 412 3.1.5 **Pressure** 413 In almost all the sites with available data (apart from the Spanish), the NPF probability presented a 414 positive relation with high significance at all types of sites. The greater significance found at the rural sites indicates the increased importance of meteorological conditions in the occurrence of NPF 415





416 events at this type of site. The growth rate also presented a similar picture, with positive relations at all the background sites of this study except the ones in Greece and FINUB (though with low R² at 417 0.02). This is probably associated with the seasonal variation found in Greece where higher growth 418 419 rates were found in summer, a period when increased wind speeds and lower atmospheric pressure 420 was found due to the Etesians (Kalkavouras et al., 2017). An interesting find is the negative slopes found at all the roadsides, though the significance of these results is relatively low ($R^2 < 0.43$) and 421 422 always lower compared to the rural sites. The effects of pressure above are not likely to be 423 important. Once again however, this is not an independent variable and higher pressure in summer 424 tends to be associated with higher insolation and temperatures and lower RH. Since most events 425 occur in the warmer months of the year, this is probably the explanation for the apparent effects of pressure. The formation rate presented relations of low significance for the sites of this study. Due 426 to this, pressure should not be an important factor for the formation rate at any type of site. 427 428

The normalised slopes did not present any clear trends, even for the NPF probability for which the results presented significant relations at almost all sites.

3.2 Atmospheric Composition

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The slopes and R² from the analysis of a number of air pollutants and the condensation sink, as well as the average conditions of these variables are found in Table 4. The results for each site and variable are found in Figure S1.



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3.2.1 Sulphur dioxide (SO₂)

Sulphur dioxide is considered as one of the main components that participate in the NPF process. According to nucleation theories and observations, H₂SO₄ is the most important compound from which the initial clusters are formed, as well as one of the candidate compounds for the initial steps of particle growth (Kirkby et al., 2011; Nieminen et al., 2010; Sipila et al., 2010). As H₂SO₄ in the atmosphere is produced from oxidation reactions of SO₂ it would be expected that increased concentrations of the latter would be associated with increased values for all the variables associated with the NPF process. Contrary to this though, the relation of SO₂ concentrations with NPF probability was found to be negative at all the sites in this study with available data. This relation was relatively strong ($R^2 > 0.50$) in most areas with an increased significance at roadsides compared to their respective rural sites. As this is a negative relation, this may indicate that SO₂ is in sufficient concentrations for H₂SO₄ formation, thus not suppressing the occurrence of NPF events, as well as showing that in increased concentrations, it is a more important factor (or surrogate for a factor) in preventing the occurrence of NPF events within the urban environment, as probably higher SO₂ is associated with increased co-emitted particle pollution and hence CS. The growth rate on the other hand, presented mixed results and the significance of the relationships is low in most cases, which makes these results untrustworthy. Finally, the relation of SO₂ concentrations with the formation rate was found to be positive at all sites but SPARU and FINRU (which had the lowest concentrations across the sites of this study). The significance of this relationship was rather low for all but the roadsides. This suggests that higher H₂SO₄ concentrations



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promoting nucleation itself because of the competing effect of condensation onto the pre-existing 457 particle population. 458 459 The normalised slopes a_N^* were found to be more negative at the background sites compared to 460 their respective roadsides, as well as being less negative in the UK (where SO₂ is in greater 461 462 abundance) compared to the other sites with relatively significant relations. Plotting the average SO₂ concentrations with the normalised slopes a_N^* for the all sites (though not all had significant 463 relations), a positive relation with relatively high R² (when the extreme values from Marylebone 464 Road-UKRO are removed) is found which might indicate that while increased concentrations are a 465 negative factor in NPF event occurrence at a given site, in general the sites with higher SO₂ 466 concentrations on average present higher probability for NPF events (Figures 7a and 7b). This 467 appears to be in agreement with Dall'Osto et al. (2018) who discussed the variable role of SO₂ 468 depending on its concentrations. No significant relations were found for the values of a_J* as in most 469 470 cases these relations were rather weak. 471 472 3.2.2 Nitrogen oxides or nitrogen dioxide (NO_x or NO₂) 473 NO_x and NO₂ are directly associated with pollution, which can be a limiting factor for NPF events

favour increased formation rates (i.e. more particles can be formed), rather than necessarily

as it increases the CS and may suppress the events (An et al., 2015), though with the reduction of

SO₂ concentrations achieved the last couple of decades, there is possibility for oxidation products



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of NO_x to become an important component for NPF (Wang et al., 2020). For almost all sites (apart from GRERU) with available data a negative relation between the NPF probability and NO_x (or NO₂) concentrations (depending on what data was available) was found. Similarly, for all the sites but SPARU and GRERU, the correlations were strong with $R^2 > 0.43$. The rural background sites had a weaker relation between the two variables compared to the urban sites, which is probably associated with them having rather low concentrations of NO_x (or NO₂) and variability, making the variations of this factor less important. Growth rate had weaker correlations with NO_x and different trends between the sites, either being positive or negative. The variable effect of NO_X on particle growth, shifting HOMs' volatility, was previously discussed by Yan et al. (2020). While variability was found for the background sites, all roadsides regardless of the strength of the relation had positive relation between NO_x and the growth rate. This may indicate the different components associated with the growth process at each type of site which, as found in other studies can be related to compounds associated with combustion processes that take place within the urban environment (Guo et al., 2020; Wang et al., 2017a). The formation rate presents few cases of strong relations, with variable trends (positive and negative). While much effort was made to isolate the effect of NPF events by taking a shorter time frame before the event, the effect of local pollution is still included, especially at the urban sites.

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The normalised slopes do not provide a significant result for the relationship of this variable with either the probability of the events or the formation rate. The only noteworthy points are the more





negative a_N^* at the rural background sites compared to the roadsides in all the areas studied, which shows the increased importance of a clean environment for NPF events to occur in areas where condensable compounds are in lesser abundance, such as a rural environment. Additionally, the negative slopes found at all the roadside sites, which increases the confidence that the events extracted at the roadsides are not pollution incidents but NPF events. However, it appears that traffic pollution favours higher particle growth rates, although the components responsible for this effect are unknown.

3.2.3 Ozone (O₃)

Ozone is typically the result of atmospheric photochemistry and is itself a source of hydroxyl radical through photolysis, or ozonolysis of alkenes both during daytime and night-time (Fenske et al., 2000). It might therefore be expected to act as an indicator of photochemical activity which promotes the oxidation of SO₂ and VOCs. Ozone concentrations may be directly related to the solar radiation intensity as well as the pollution levels in the area studied, and O₃ is considered as a positive factor in the occurrence of NPF events (Woo et al., 2001; Berndt et al., 2006). As for the solar radiation, there is a strong relation between O₃ concentration and the probability for NPF events. This positive relation was found to be stronger for the sites in northern Europe, while it was not significant for the sites from southern Europe (Spanish sites and GRERU), possibly indicating that O₃ is a less important factor at the southern sites. Specifically for the Spanish sites which have the highest average concentrations of O₃ with some extreme values (Querol et al., 2017), the

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relation of O₃ concentrations with the NPF probability presents a unique trend, having a clear peak then a steady decline at both sites (though at different O₃ concentrations), which is also responsible for the low correlations found (this trend seems to also occur at SPARU for the growth rate and to a lesser extent for the formation rate as well, though for different O₃ concentration ranges). The specific variability found at the Spanish sites was also studied by Carnerero et al., (2019). For sites with a marked seasonal variation in ozone, associations with NPF may be artefactual due to correlations with other variables such as temperature, RH and solar radiation.

Unlike the solar radiation though, the growth rate presents a negative relation at the sites where the relation between these two variables was significant (UKRU, UKUB, DENUB and FINRU), which

relation between these two variables was significant (UKRU, UKUB, DENUB and FINRU), which might either be an indication of a polluted background that may have a negative effect in the growth of the newly formed particles (though the trends found for NO_x indicate differently) or specific chemical processes which cannot be identified due to the lack of detailed chemical composition data. A significant relation between O₃ and the formation rate was only found for a few sites (though the trends become a lot clearer if some values are removed from the extreme lower or higher end). This way the relations become strong, but positive, for some areas and negative for some others without any clear trend (type or location of the site, O₃ concentrations etc.). No clear relation between these two variables was found as the sites with strong relation have both positive and negative relationships and as a result no confident conclusions can be drawn.





As the correlations found were strong the normalised slopes for NPF probability, when plotted against the average concentrations of O_3 , present a negative correlation with relatively high R^2 (0.64), indicating that the O_3 is a more important factor in the occurrence of NPF events when in lower concentrations (Figure 8). Finally, though with a low level of confidence for the southern sites, the a_N^* were smaller at the southern sites compared to those in the north, up to one order of magnitude between the FINRU (furthest north rural background) and GRERU (furthest south rural background).

3.2.4 Organic compounds

544 3.2.4.1 Particulate organic carbon (OC)

Organic carbon (OC) compounds are considered as components with importance in the growth of newly formed particles, with a role that becomes increasingly important as the size of the particles becomes larger (Nieminen et al., 2010; Zhang et al., 2012; Shrivastava et al., 2017). Particulate OC, the data for which are available in the present study, can be associated with pollution, especially in the urban environment. Only a few of the sites of the present study were found to have a strong negative relationship ($R^2 > 0.50$) of particulate OC with the NPF probability (UKUB, UKRO and DENRU). Regardless though of the strength of this relation, all other sites (apart from FINRU) had a negative relationship between these two variables as well, consistent with increased concentrations of particulate OC being associated with increased pollution, which is a suppressing factor in the occurrence of NPF events. Growth rate on the other hand was found to have a slight





positive relation (R² > 0.40) for most of the sites. This relation appeared to be stronger (higher R²) at the roadsides with available data compared to their respective rural background sites. The relation between particulate OC and the growth rate was positive at all the sites with available data regardless of their significance showing that, despite its effect in the occurrence of NPF events, it is still a favourable variable for the growth of the particles. The formation rate was found to have a significant relation with particulate OC concentrations at half of the sites with available data (UKUB, UKRO, DENRU, DENRO).

The normalised slopes for this variable did not present any noteworthy relations with either the type of site or the concentrations of OC at a given site.

3.2.4.2 Volatile organic compounds (VOCs)

Many volatile organic compounds have been found to be associated with the NPF process. Benzene, toluene, ethylbenzene, m+p-xylene, o-xylene and trimethylbenzenes have been reported to be able to form Highly Oxygenated Organic Molecules (HOMs) in flow tubes (Wang et al., 2017a; Molteni et al., 2018), which may act as contributors to particle nucleation and/or growth. Xylenes, and to a lesser extent trimethylbenzenes, are the most efficient at forming HOMs. Benzene and toluene are less efficient and will form more volatile HOMs. These HOMs may all be too volatile to form new particles, though this is not yet confirmed. Chamber studies involving H₂SO₄ and trimethylbenzene oxidation products were associated with high formation rates when measuring J_{1.5} (Metzger et al.,





575 2010). All these HOMs though will be sufficiently involatile to contribute to particle growth. Those 576 with higher oxygen content or carbon number will be classed as LVOC and if they dimerise, they will form ELVOC (Bianchi et al., 2019). Monoterpenes can also form HOMs which drive both the 577 578 formation (Ehn et al., 2014; Riccobono et al., 2014) and growth (Tröstl et al., 2016), while isoprene 579 can act as a sink for hydroxyl radical (Kiendler-Scharr et al., 2009) and is not as effective in HOM 580 and secondary organic aerosol formation compared to monoterpenes (McFiggans et al., 2019). 581 582 Volatile organic compound data were available for three of the sites of this study (Table S2). Two 583 of the sites with VOC data were from the rural background and the roadside in the UK. Most of the 584 compounds are associated with combustion sources and were found to have a negative relationship with NPF event occurrence at both sites, with high R² in most cases. Additionally, isoprene, which 585 may have either biogenic or anthropogenic sources (Wagner and Kuttler, 2014) was also found to 586 587 have a negative relationship with NPF event occurrence at Marylebone Road-UKRO, though with low R². This result is in line with the VOCs being strongly correlated with particulate OC (which 588 589 presented a negative relation with NPF event probability, as discussed in Section 3.2.4.1), as well as 590 with the CS (which also presented a negative relationship with NPF event probability, as mentioned 591 in Section 3.2.6), further associating these compounds with combustion emissions. 592 593 Growth rate was found to have a positive relationship with VOCs in almost all cases for both UK 594 sites. Few exceptions were found (with only 1,3 butadiene having a relatively high R²) which



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presented a negative relationship with the growth rate in rural Harwell-UKRU. Finally, the formation rate presented a different behaviour between the two sites. At Harwell-UKRU, the relationship was unclear in most cases, with a group of VOCs presenting a negative relationship with the formation rate (ethane, ethene, propane, 1,3 butadiene, toluene, ethylbenzene, o-xylene and 1,2,4 trimethylbenzene – with $R^2 > 0.40$), two VOCs presented a rather clear positive relationship with the formation rate (iso-pentane and 2-methylbenzene) and the rest of the VOCs had an unclear relationship. At Marylebone Road-UKRO though, VOCs presented a positive relationship with the formation rate (for particles of diameter 16 nm). This is probably due to the fact that these VOCs are associated with pollution emissions (as mentioned earlier) and though a smaller time window was chosen to avoid including the effect of the morning rush hour traffic, this is very difficult in the traffic polluted environment of Marylebone Road-UKRO. As Hyvtiälä (FINRU) is a rural background site far from the direct effect of combustion emissions, different VOCs were measured, which mainly originate from biogenic sources rather than anthropogenic ones. The results were mixed and less clear compared to those from the UK sites (mainly due to the smaller dataset), and three groups were found depending on their relationship with NPF probability. The first group, including acetonitrile, acetic acid and Methyl Ethyl Ketone (MEK) presented a slight positive relation. The second group presented a negative relation, with the VOCs in this group being MEK, monoterpenes, benzene, isoprene and toluene (only the last two

have $R^2 > 0.50$). Finally, the third group included VOCs that presented a peak and then a decline for





higher concentrations including methanol, and acetone. Two groups of VOCs were found depending on their relationship with the growth rate. The ones with a positive relation being methanol, acetonitrile, acetone, acetic acid, isoprene, MEK, monoterpenes and toluene, while acetaldehyde, MEK and benzene had a negative relationship, with relatively high R² in most cases. Finally, the results with the formation rate were unclear with only a handful presenting weak positive (methanol, acetic acid and benzene) or negative (MEK) relations that do not appear to be significant. The normalised slopes cannot be used for VOCs as there are very few sites with available data.

3.2.5 Sulphate (SO₄-2)

Sulphate (SO₄²⁻) is a major secondary constituent of aerosols. Secondary SO₄²⁻ aerosols largely arise from either gas phase reaction between SO₂ and OH, or in the aqueous phase by the reaction of SO₂ and O₃ or H₂O₂, or NO₂ (Hidy et al., 1994). In environments where SO₄²⁻ chemistry is dominant (i.e. remote areas), SO₄²⁻ and ammonium (bi) sulphate ((NH₄)₂SO₄ and NH₄HSO₄) particles are a large relative contributor to aerosol mass, while this contribution is lower in environments where other emissions are also significant (i.e. urban areas where the secondary NO₃⁻ relative contribution is a lot higher). While not well established, a possible relation of SO₄²⁻-containing compounds and variables of NPF events was found in previous studies (Beddows et al., 2015; Minguillón et al., 2015; Wang et al., 2017b). In the present study, only a few sites had SO₄²⁻ data available, for PM₁ (FINRU), PM_{2.5} (Danish sites) or PM₁₀ (rest of the sites). While this data cannot be considered as





635 directly associated with the ultrafine particles, for two sites with available AMS data for ultrafine particles, the direct comparison between SO₄²- aerosol in PM and in the range of particles of about 636 50 nm, very high correlations were found (results not included). For all the sites with available data 637 638 the NPF probability presented a negative relation. The significance of this relations was found to be relatively high ($R^2 > 0.50$) only for background sites (apart from GERRU, which has rather low 639 concentrations and probably different mechanisms for the NPF events). Similarly, the growth rate 640 presented a more significant relation ($R^2 > 0.40$) for the same background sites (apart from FINRU). 641 642 though this relationship was found to be positive at all sites regardless of its significance. Finally, 643 the formation rate did not present a clear trend as it was found to have both negative and positive relations for different sites. This relation was significant only for two rural sites (UKRU and 644 645 DENRU) and as a result no assumptions can be made.

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The normalised slopes cannot be used for any analysis on sulphate as the measurements available are from different particle sizes.

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3.2.6 Gaseous ammonia (NH₃)

Ammonia (NH₃) can be an important compound in the nucleation process according to the ternary theory (Napari et al., 2002). It was found that elevations in NH₃ concentrations can lead to elevations to NPF rate (Lehtipalo et al., 2018) and it was also found to be an important factor for NPF event occurrence even when stronger bases are present in high concentrations (Glasoe et al.,





2015). No significant variation was found though between event and non-event days in a previous study in Harwell-UKRU (Bousiotis et al., 2019). Data for gaseous ammonia were only available for Harwell-UKRU and presented a positive relation with NPF probability, until reaching a peak point. Further increase in NH₃ concentrations presented a decline with NPF probability, which might be due to its association with increased pollution levels. Interesting though is that it presented a clear positive relation with both the growth rate (though it also appears to decline at high concentrations) and the formation rate.

3.2.7 Condensation sink (CS)

The CS is a measure of the rate at which molecules will condense onto pre-existing aerosols (Lehtinen et al., 2003). It is highly dependent on the number and size of the particles in the atmosphere and as a result it is expected to be affected by both the local emissions within the urban environment as well as the formation and growth of the particles due to NPF events. As a result, for the specific metric a time frame before the events are in full development was chosen (05:00 to 10:00 LT) to avoid including the effect of the NPF events and provide a picture of the atmospheric conditions that preceded the NPF events. With this data, the NPF probability presented very strong relations with the condensation sink. Two groups of sites were found though; those which had a positive relation and those with a negative relation. In the first group are the sites in Germany and Greece while all others had a negative relation. This grouping follows the trend between the countries, the sites of which presented a greater (the ones with the positive slopes) or smaller CS on

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675 NPF event days, though it is unknown what causes this behaviour (at the German sites and GREUB it may be associated with the very high formation rates on NPF event days). While the slopes from 676 this analysis cannot be used for direct comparisons, a trend was found for which the slopes were 677 678 more positive or negative at the rural sites compared to their respective roadsides, which might 679 indicate the greater importance of the variability of the CS at the rural sites in the occurrence of NPF events. 680 681 The growth rate was positively correlated with the CS for most of the sites, with strong relations 682 (high R²) for about half of them. As the CS is a metric of pre-existing particles, it is also associated 683 684 with the level of pollution in a given area. The increased significance and slope found at the rural sites probably indicates the importance of enhanced presence of condensable compounds in a 685 cleaner environment, which in many cases are associated with the moderate presence of pollution. 686 687 The formation rate was also found to have a positive relation with the CS. This relation was more significant at the roadsides of this study, a result which to some extent is biased by the presence of 688 689 increased traffic emissions found in the timeframe chosen. While to an extent, increased presence of 690 condensable compounds can be favourable for greater formation rates, this result should be 691 considered with great caution. 692 The normalised slopes a_N^* followed a similar trend as those found with the initial analysis. These 693 694 slopes were found to be more positive or negative, depending on the trend of the given area, at the



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rural sites compared to their roadsides. The urban background sites did not always have a uniform behaviour (though in UK, Denmark and Finland these were between the rural site and the roadside), due to their more diverse character compared to the other two types of sites.

3.3 Association of the Effect of the Variables

The Pearson correlation coefficients for the variables studied on each site are found in Table S1. The relatively strong relation between the solar radiation, temperature and O₃ found, as well as their anticorrelation with the RH may lead to the conclusion that not all these factors play a role in NPF events, but their visible effect is the result of their relationship with each other. There is a similar case with the association of the CS and NO_x (or NO₂), and OC, as well as SO₂, especially at urban sites. However, the factors affect different outcomes differently, as for example the solar radiation intensity does not seem to be as important a factor for the growth rate as temperature, or O₃ does not seem to be strongly associated with either the formation or the growth rate. This is further established by the fact that some of these variables do not correlate well at the southern sites, but still appear to be associated with either the probability of NPF events or the growth or nucleation rate. The effects of all of these factors have been demonstrated in both laboratory and atmospheric studies in the past and were discussed earlier in this paper. By the analysis provided in the present study, the effect of each of these variables is further established, providing an association of each one of these variables with either the formation or the growth mechanism. However, RH does not seem to be a consistent factor in any mechanism, and it appears that its effect is dependent on





location specific conditions, although it was the variable with the most consistent relation with NPFevent probability at almost all sites.

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3.4 Relationship to a previous multi-station European study

The findings of our study in respect of the background sites show many similarities with the conclusions drawn in the previous multi-station study in Europe by Dall'Osto et al. (2018) despite the two studies using several different sampling stations as well as some in common. Both studies point towards the influence of variables such as solar radiation and CS upon the occurrence of NPF events. The previous study suggested that different compounds participate in the growth of the particles, depending on the area considered. Thus, for northern and southern sites the growth of the particles is suggested to be driven mainly by organic compounds, while for the sites in central Europe sulphate plays a more important role. These findings are confirmed by the present study, as the growth rate was found to correlate better with organic compounds for the rural sites in Finland and Greece, while SO_4^{2-} presented a stronger relation with the growth rate for the Danish and German sites (the latter presented high slope values but low R² due to a decline at higher SO₄²concentrations, probably associated with NPF events being suppressed by increased pollution). The growth of the particles at the rural background site in the UK, characterised as "Overlap" in the previous study, was found to be strongly associated with both organic compounds and sulphate, consistent with it being in the central group.





735 The seasonality of NPF events at northern sites was hard to explain in the previous study, and the 736 possible effect of low temperature was considered. In the present study, the Finnish background 737 sites presented a double-peak relation of NPF probability with temperature, with one of the peaks 738 being below zero degrees. This might point to the possibility of different compounds driving the 739 events for different temperature ranges, as well as the increased nucleation rate of H₂SO₄ at lower 740 temperatures (Kirkby et al., 2011; Yan et al., 2018), which makes the occurrence of NPF events 741 more probable at lower temperatures in a region with low SO₂ concentrations.

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743 4. CONCLUSIONS

More than 85 site-years of data from 16 sites from six countries in Europe were analysed for NPF events. A total of 1950 NPF events with consequent growth of the newly formed particles were extracted and with the use of binned linear regression, the relation between three variables associated with NPF events (NPF event probability, formation and growth rate) with meteorological conditions and atmospheric composition was studied. Among the meteorological conditions, solar radiation, temperature and atmospheric pressure presented a positive relation with NPF event occurrence, and either promoting the formation or growth rate. Relative humidity presented a negative relation with NPF event probability which in most cases was associated with it being a limiting factor on particle formation at higher values. Wind speed on the other hand presented variable results, appearing to depend on the location of the sites rather than their type. This shows that while wind speed can be a factor in NPF event occurrence, the origin of the incoming air





755 masses also plays a very important role. In most cases, meteorological conditions appeared to be 756 more important factors in NPF event occurrence at rural sites compared to urban sites, suggesting that NPF events are driven more by them at this type of site. Additionally, while some 757 758 meteorological variables appeared to play a crucial role in the occurrence of NPF events, this role 759 appears to become less important at higher values when a positive relation was found (or lower when a negative relation was found). 760 761 The results for the levels of atmospheric pollutants presented a more interesting picture as most of 762 763 these, which appear to be either directly or indirectly associated with the NPF process were found to have negative relations with NPF probability. This is probably due to the fact that increased 764 concentrations of such compounds are associated with more polluted conditions, which are a 765 limiting factor in the occurrence of NPF events, as was found with the negative relation between the 766 CS and NPF probability in most cases. Thus, SO₂, NO_x (or NO₂), particulate OC and SO₄² 767 concentrations were negatively correlated with NPF probability in most cases. Average SO₂ 768 769 concentrations though appeared to correlate positively with the normalised NPF event probability slopes with relatively significant correlation, indicating that while increasing concentrations have a 770 771 negative impact in the occurrence of NPF events at a given site, in general sites with higher SO₂ 772 concentrations have higher probability for NPF events. On the other hand though, these compounds 773 in many cases had a positive relation (not always though with high significance) with the other variables considered. Thus, particulate OC (and VOCs where data were available) and SO₄²-774





consistently had a positive relation with the growth rate, while SO₂ was positively associated with both the formation and growth rate in most cases. Finally, O₃ was positively correlated with NPF event probability at all sites in this study, though it presented variable results with the other two variables. As with some meteorological conditions it was found that at sites with increased concentrations of O₃, its importance as a factor was decreased, which to an extent can be related with high CS associated with peak summer O₃ days in southern Europe.

The present study attempts to explain the effect of several meteorological and atmospheric variables on the occurrence and development of NPF events, by using a large-scale dataset. It should be noted that the variables considered are in many cases inter-related (e.g. temperature and RH) and this complicates considerably the interpretation in terms of causal factors. Large datasets are very useful in providing with more uniform results by removing the possible bias of short period extremities, which may lead to wrong assumptions. Following from this, the importance of a high-resolution measurement network, both site and timewise is underlined, as it can help in elucidating the mechanisms of new particle formation in the real atmosphere.

DATA ACCESSIBILITY

- 792 Data supporting this publication are openly available from the UBIRA eData repository at
- 793 https://doi.org/https://doi.org/10.25500/edata.bham.00000491





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| 1377 | | r |
| 1378 | Figure 7b: | Relation of average SO ₂ concentrations and normalised slopes a_N^* for the sites of the |
| 1379 | C | present study (UKRO not included). |
| | | |



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Relation of average O_3 concentrations and normalised slopes $a_N^{\ *}$ for the sites of the 1380 **Figure 8:**

present study.





Table 1: Location and data availability of the sites.

| Table | 1: Location and data avail | l | Meteorological | Data | 1 |
|-------|---|---|------------------------------------|------------------------------|------------------------------------|
| Site | Location | Available data | data location | availability | Reference |
| UKRU | Harwell Science Centre, Oxford, 80 km W of London, UK (510 34' 15" N; | SMPS (16.6 - 604 nm, 76.5% availability), | On site | 2009 - 2015 | Charron et al., |
| UKUB | lo 19' 31" W) North Kensington, 4 km W of London city centre, UK (510 31' 15" N; 00 12' 48" W) | ammonia SMPS (16.6 - 604 nm, 83.3% availability), NOx, SO2, O3, OC, SO42- | Heathrow airport | 2009 - 2015 | 2013 Bigi and Harrison, 2010 |
| UKRO | Marylebone Road, London, UK (510 31' 21" N; 00 9' 16" W) | SMPS (16.6 - 604 nm, 74.3% availability), NOx, SO2, O3, OC, SO42- | Heathrow airport | 2009 - 2015 | Charron and Harrison, 2003 |
| DENRU | Lille Valby, 25 km W of Copenhagen, (550 41' 41" N; 120 7' 7" E) (2008 – 6/2010) Risø, 7 km north of Lille Valby, (55° 38' 40" N; 12° 5' 19" E) (7/2010 – 2017) | DMPS and CPC (5.8 - 700 nm, 68.3% availability), NOx, SO2, O3, OC, SO42- | H.C. Ørsted – Institute station | 2008 – 2017 | Ketzel et al., 2004 |
| DENUB | H.C. Ørsted – Institute, 2 km NE of the city centre, Copenhagen, Denmark (550 42' 1" N; 120 33' 41" E) | availability), | On site | 2008 – 2017 | Wang et al., 2010 |
| DENRO | H.C. Andersens Boulevard, Copenhagen, Denmark (55o 40' 28" N; 12o 34' 16" E) | DMPS and CPC (5.8 - 700 nm, 65.7% availability), NOx, SO2, O3, OC, SO42- | H.C. Ørsted – Institute station | 2008 – 2017 | Wang et al., 2010 |
| GERRU | Melpitz, 40 km NE of Leipzig, Germany (510 31' 31.85" N; 120 26' 40.30" E) | TDMPS with CPC (4.8 - 800 nm, 87.2% availability), OC, SO42- | On site | 2008 – 2011 | Engler et al., 2007 |
| GERUB | Tropos, 3 km NE from the city centre of Leipzig, Germany (510 21' 9.1" N; 120 26' 5.1" E) | TDMPS with CPC (3 - 800 nm, 90.4% availability) | On site | 2008 – 2011 | Costabile et al., 2009 |
| GERRO | Eisenbahnstraße, Leipzig, Germany (510 20' 43.80" N; 120 24' 28.35" E) | TDMPS with CPC (4 - 800 nm, 68.3% availability) | Tropos station | 2008 – 2011 | Birmili et al., 2016 |
| FINRU | Hyytiälä, 250 km N of Helsinki, Finland (610 50' 50.70" N; 240 17' 41.20" E) | TDMPS with CPC (3 – 1000 nm, 98.2% availability), NOx, SO2, O3, VOCs | On site | 2008 – 2011 & 2015 – 2018 | Aalto et al., 2001 |
| FINUB | Kumpula Campus 4 km N of the city centre, Helsinki, Finland (60o 12' 10.52" N; 24o 57' 40.20" E) | TDMPS with CPC (3.4 - 1000 nm, 99.7% availability) | On site | 2008 – 2011 & 2015 – 2018 | Järvi et al., 2009 |
| FINRO | Mäkelänkatu street, Helsinki, Finland (60o 11' 47.57" N; 24o 57' 6.01" E) | DMPS (6 - 800 nm, 90.0% availability), NOx, O3 | Pasila station and on site | 2015 – 2018 | Hietikko et al., 2018 |
| SPARU | Montseny, 50 km NNE from Barcelona, Spain (41o 46' 45" N; 2o 21' 29" E) | SMPS (9 – 856 nm, 53.7% availability), NO2, SO2, O3 | On site | 2012 - 2015 | Dall'Osto et al., 2013 |
| SPAUB | Palau Reial, Barcelona, Spain (41o 23' 14" N; 2o 6' 56" E) | SMPS (11 – 359 nm, 88.1% availability), NO2, SO2, O3 | On site | 2012 – 2015 | Dall'Osto et al., 2012 |
| GRERU | Finokalia, 70 km E of Heraklion, Greece (35o 20' 16.8" N; 25o 40' 8.4" E) | SMPS (8.77 - 849 nm, 85.0% availability), NO2, O3, OC | On site | 2012 – 2018 | Kalkavouras et al., 2017 |





| GREUB | "Demokritos", 12 km NE from the city centre, Athens, Greece (370 59' 41.96" N; 230 48' 57.56" E) | SMPS (10 – 550 nm, 88.0% availability) | On site | Mølgaard et al., 2013 |
|-------|--|--|---------|--------------------------|





Table 2: Frequency and formation rate of NPF events for the sites of the study.

| | Frequency of | J_{10} |
|-------|----------------|---------------------------------------|
| Site | NPF events (%) | (N cm ⁻³ s ⁻¹) |
| UKRU | 7.0 | 8.69E-03* |
| UKUB | 7.0 | 1.42E-02* |
| UKRO | 6.1 | 3.75E-02* |
| DENRU | 7.9 | 2.57E-02 |
| DENUB | 5.8 | 2.40E-02 |
| DENRO | 5.4 | 8.07E-02 |
| GERRU | 17.1 | 9.18E-02 |
| GERUB | 17.5 | 1.02E-01 |
| GERRO | 9.0 | 1.38E-01 |
| FINRU | 8.7 | 1.19E-02 |
| FINUB | 5.0 | 2.49E-02 |
| FINRO | 5.1 | 6.94E-02 |
| SPARU | 12 | 1.54E-02 |
| SPAUB | 13.1 | 2.12E-02 |
| GRERU | 6.5 | 4.90E-03 |
| GREUB | 8.5 | 4.41E-02 |

* J₁₆ calculated





Table 3: Normalised slopes (non-normalised for growth rate), R^2 and p-values (- for values >0.05) for the relation between meteorological conditions and NPF event variables.

| | 1 octive con more | | | l shortwave | | | | | | |
|-------|---------------------|----------------|---------|---------------------------|----------------|---------|-----------------------|----------------|---------|---------|
| Site | $a_N* (W^{-1} m^2)$ | \mathbb{R}^2 | р | \mathbf{a}_{G} | \mathbb{R}^2 | p | $a_{J}*(W^{-1}m^{2})$ | \mathbb{R}^2 | р | Average |
| UKRU* | 1.21E-03 | 0.94 | < 0.001 | 6.53E-05 | 0.11 | ı | 6.28E-04 | 0.93 | < 0.001 | 443 |
| UKUB* | 6.81E-04 | 0.90 | < 0.001 | -8.26E-05 | 0.10 | ı | 1.49E-04 | 0.19 | - | 448 |
| UKRO* | 8.69E-04 | 0.98 | < 0.001 | -7.75E-06 | 0.00 | ı | 2.66E-04 | 0.64 | < 0.005 | 464 |
| DENRU | 2.22E-03 | 0.88 | < 0.001 | 4.24E-04 | 0.20 | ı | 1.38E-03 | 0.64 | < 0.001 | 115 |
| DENUB | 1.87E-03 | 0.91 | < 0.001 | 1.47E-04 | 0.03 | - | 8.98E-04 | 0.48 | < 0.01 | 115 |
| DENRO | 2.46E-03 | 0.95 | < 0.001 | 1.27E-04 | 0.01 | - | 6.77E-04 | 0.50 | < 0.005 | 117 |
| GERRU | 2.87E-03 | 0.98 | < 0.001 | 9.88E-04 | 0.72 | < 0.01 | 1.45E-03 | 0.81 | < 0.001 | 130 |
| GERUB | 3.18E-03 | 0.97 | < 0.001 | 7.28E-04 | 0.51 | < 0.005 | 1.53E-03 | 0.69 | < 0.001 | 114 |
| GERRO | 2.40E-03 | 0.95 | < 0.001 | -5.89E-04 | 0.09 | - | 9.95E-04 | 0.59 | < 0.005 | 114 |
| FINRU | 2.63E-03 | 0.76 | < 0.001 | 1.01E-03 | 0.57 | < 0.01 | 2.04E-03 | 0.82 | < 0.001 | 91.5 |
| FINUB | 1.38E-03 | 0.37 | - | 1.81E-04 | 0.08 | ı | 8.99E-04 | 0.25 | - | 111 |
| FINRO | 1.76E-03 | 0.59 | < 0.005 | 9.15E-04 | 0.34 | < 0.005 | 4.45E-04 | 0.03 | - | 114 |
| SPARU | 3.46E-04 | 0.35 | < 0.05 | 5.68E-04 | 0.13 | ı | 1.97E-03 | 0.74 | < 0.001 | 162 |
| SPAUB | 5.92E-04 | 0.58 | < 0.05 | 6.98E-04 | 0.23 | - | 1.58E-03 | 0.81 | < 0.001 | 180 |
| GRERU | 4.10E-04 | 0.52 | < 0.001 | 7.14E-04 | 0.55 | < 0.001 | -6.30E-04 | 0.05 | - | 201 |
| GREUB | 3.49E-04 | 0.31 | - | -1.10E-04 | 0.02 | - | 8.97E-04 | 0.34 | < 0.05 | 183 |

^{*} Global solar irradiation measurements in kJ m⁻²

| | | | | Relative H | umidit | y (%) | | | | |
|-------|-------------------------------------|----------------|---------|---------------------------|----------------|---------|-------------------------------------|----------------|---------|---------|
| Site | a _N * (% ⁻¹) | \mathbb{R}^2 | р | \mathbf{a}_{G} | R ² | р | a _J * (% ⁻¹) | \mathbb{R}^2 | р | Average |
| UKRU | -5.89E-02 | 0.85 | < 0.001 | 1.69E-03 | 0.02 | - | -3.35E-02 | 0.85 | < 0.001 | 79.7 |
| UKUB | -3.42E-02 | 0.94 | < 0.001 | 8.23E-03 | 0.24 | - | -5.66E-03 | 0.19 | - | 75.3 |
| UKRO | -5.09E-02 | 0.85 | < 0.001 | 7.03E-03 | 0.25 | - | -1.49E-02 | 0.46 | < 0.05 | 74.5 |
| DENRU | -3.90E-02 | 0.95 | < 0.001 | 9.42E-03 | 0.74 | < 0.001 | 5.45E-04 | 0.00 | - | 75.7 |
| DENUB | -3.14E-02 | 0.94 | < 0.001 | 3.64E-03 | 0.06 | - | 2.57E-03 | 0.00 | - | 75.7 |
| DENRO | -3.64E-02 | 0.95 | < 0.001 | -1.21E-02 | 0.22 | - | -3.91E-03 | 0.10 | - | 75.7 |
| GERRU | -5.08E-02 | 0.88 | < 0.001 | -1.30E-02 | 0.72 | < 0.001 | -2.46E-02 | 0.91 | < 0.001 | 81.9 |
| GERUB | -5.35E-02 | 0.86 | < 0.001 | -6.34E-03 | 0.67 | < 0.001 | -2.25E-02 | 0.86 | < 0.001 | 78.7 |
| GERRO | -2.83E-02 | 0.90 | < 0.001 | 3.98E-03 | 0.05 | - | -1.72E-02 | 0.81 | < 0.001 | 78.7 |
| FINRU | -4.48E-02 | 0.94 | < 0.001 | -7.07E-03 | 0.65 | < 0.001 | -2.16E-02 | 0.87 | < 0.001 | 80.1 |
| FINUB | -5.89E-02 | 0.95 | < 0.001 | 1.04E-02 | 0.26 | - | -6.52E-03 | 0.18 | - | 76.5 |
| FINRO | -3.34E-02 | 0.92 | < 0.001 | -1.47E-03 | 0.01 | - | 7.39E-03 | 0.10 | - | 71.1 |
| SPARU | -1.54E-02 | 0.90 | < 0.001 | -4.67E-03 | 0.08 | - | -7.12E-03 | 0.14 | - | 66.4 |
| SPAUB | -4.84E-02 | 0.93 | < 0.001 | 2.43E+02 | 0.50 | < 0.01 | -9.83E-03 | 0.19 | - | 69.2 |
| GRERU | -7.72E-03 | 0.22 | - | 1.06E-02 | 0.06 | - | -1.83E-01 | 0.15 | - | 70.0 |
| GREUB | -1.42E-02 | 0.62 | < 0.001 | 2.83E-03 | 0.06 | - | 4.85E-04 | 0.00 | - | 60.5 |





| Temperature (°C) | | | | | | | | | | | | |
|------------------|-------------------------|----------------|---------|---------------------------|----------------|---------|-------------------------|----------------|---------|---------|--|--|
| Site | a _N * (°C-1) | \mathbb{R}^2 | p | \mathbf{a}_{G} | \mathbb{R}^2 | p | a _J * (°C-1) | \mathbb{R}^2 | p | Average | | |
| UKRU | 1.10E-01 | 0.93 | < 0.001 | 7.85E-02 | 0.94 | < 0.001 | 8.72E-02 | 0.84 | < 0.001 | 10.6 | | |
| UKUB | 9.04E-02 | 0.98 | < 0.001 | 1.39E-01 | 0.96 | < 0.001 | 6.34E-02 | 0.73 | < 0.005 | 11.8 | | |
| UKRO | 8.22E-02 | 0.98 | < 0.001 | 3.51E-02 | 0.52 | < 0.05 | 4.32E-02 | 0.44 | < 0.05 | 12.1 | | |
| DENRU | 6.68E-02 | 0.83 | < 0.001 | 1.54E-02 | 0.08 | - | 6.68E-02 | 0.92 | < 0.001 | 9.80 | | |
| DENUB | 2.50E-02 | 0.45 | < 0.05 | 2.40E-02 | 0.33 | - | 3.05E-02 | 0.45 | < 0.05 | 9.82 | | |
| DENRO | 6.64E-02 | 0.88 | < 0.001 | 3.51E-03 | 0.00 | - | 2.96E-02 | 0.58 | < 0.005 | 10.0 | | |
| GERRU | 7.27E-02 | 0.92 | < 0.001 | 5.65E-02 | 0.92 | < 0.001 | 5.37E-02 | 0.93 | < 0.001 | 10.3 | | |
| GERUB | 8.20E-02 | 0.93 | < 0.001 | 3.38E-02 | 0.62 | < 0.001 | 4.28E-02 | 0.54 | < 0.005 | 11.1 | | |
| GERRO | 5.08E-02 | 0.89 | < 0.001 | -3.33E-03 | 0.00 | - | 1.61E-02 | 0.11 | - | 11.1 | | |
| FINRU | -2.01E-02 | 0.17 | - | 1.13E-01 | 0.79 | < 0.001 | 4.27E-02 | 0.72 | < 0.001 | 4.79 | | |
| FINUB | -4.21E-03 | 0.00 | - | 7.42E-02 | 0.83 | < 0.001 | 1.67E-02 | 0.28 | - | 6.52 | | |
| FINRO | 6.24E-02 | 0.65 | < 0.005 | 9.28E-02 | 0.87 | < 0.001 | -1.09E-02 | 0.05 | - | 7.72 | | |
| SPARU | -2.51E-02 | 0.41 | < 0.05 | 1.23E-01 | 0.92 | < 0.001 | 9.11E-02 | 0.71 | < 0.001 | 13.9 | | |
| SPAUB | -3.43E-03 | 0.02 | - | 6.67E-02 | 0.66 | < 0.005 | 1.18E-02 | 0.08 | - | 18.2 | | |
| GRERU | -4.66E-02 | 0.75 | < 0.001 | 1.74E-01 | 0.75 | < 0.001 | -9.45E-02 | 0.47 | < 0.05 | 18.2 | | |
| GREUB | -1.00E-02 | 0.25 | - | 4.67E-02 | 0.62 | < 0.005 | -2.85E-02 | 0.20 | - | 17.6 | | |

| | | | | Wind S | peed (1 | m s ⁻¹) | | | | |
|-------|--------------------------------------|----------------|---------|---------------------------|----------------|---------------------|--------------------------------------|----------------|---------|---------|
| Site | a _N * (m ⁻¹ s) | \mathbb{R}^2 | р | $\mathbf{a}_{\mathbf{G}}$ | \mathbb{R}^2 | р | a _J * (m ⁻¹ s) | R ² | р | Average |
| UKRU | 5.72E-02 | 0.20 | - | -3.04E-02 | 0.07 | - | 6.87E-03 | 0.00 | - | 3.96 |
| UKUB | 1.72E-01 | 0.87 | < 0.001 | -1.91E-01 | 0.71 | < 0.001 | 3.56E-03 | 0.00 | - | 4.16 |
| UKRO | 6.34E-02 | 0.19 | - | 3.21E-02 | 0.02 | - | 7.28E-02 | 0.45 | < 0.005 | 4.14 |
| DENRU | 1.08E-01 | 0.88 | < 0.001 | -2.33E-01 | 0.74 | < 0.001 | 1.28E-01 | 0.44 | < 0.01 | 4.17 |
| DENUB | 1.50E-01 | 0.90 | < 0.001 | -3.33E-02 | 0.10 | - | 8.31E-02 | 0.19 | - | 4.17 |
| DENRO | 1.65E-01 | 0.89 | < 0.001 | -1.51E-01 | 0.49 | < 0.001 | 9.08E-03 | 0.00 | - | 4.16 |
| GERRU | -1.06E-01 | 0.57 | < 0.005 | -2.26E-01 | 0.83 | < 0.001 | -5.32E-03 | 0.00 | - | 2.58 |
| GERUB | -1.27E-01 | 0.52 | < 0.01 | -1.41E-01 | 0.60 | < 0.005 | -3.32E-02 | 0.04 | - | 2.33 |
| GERRO | -2.40E-01 | 0.56 | - | -2.54E-01 | 0.38 | - | -1.30E-01 | 0.22 | - | 2.33 |
| FINRU | 1.62E-01 | 0.63 | < 0.005 | -1.29E-01 | 0.16 | < 0.05 | 7.99E-02 | 0.07 | - | 1.31 |
| FINUB | -3.17E-02 | 0.08 | - | 7.26E-02 | 0.20 | < 0.05 | -9.74E-02 | 0.17 | - | 3.43 |
| FINRO | 8.62E-02 | 0.51 | < 0.05 | -1.60E-01 | 0.32 | < 0.05 | -1.86E-01 | 0.32 | - | 4.26 |
| SPARU | -2.20E-02 | 0.02 | - | 3.80E-01 | 0.31 | - | 5.74E-02 | 0.02 | - | 0.94 |
| SPAUB | 2.90E-01 | 0.93 | < 0.001 | 7.71E-02 | 0.24 | - | -5.90E-02 | 0.05 | - | 2.05 |
| GRERU | 4.37E-02 | 0.54 | < 0.001 | 1.01E-01 | 0.36 | < 0.005 | 1.73E-03 | 0.00 | - | 6.06 |
| GREUB | -1.13E-01 | 0.47 | < 0.01 | -1.88E-01 | 0.50 | < 0.005 | -3.78E-02 | 0.01 | - | 1.87 |





| | Atmospheric Pressure (mbar) | | | | | | | | | | | | |
|-------|--|----------------|---------|---------------------------|----------------|---------|--|----------------|--------|---------|--|--|--|
| Site | a _N * (mbar ⁻¹) | \mathbb{R}^2 | p | \mathbf{a}_{G} | \mathbb{R}^2 | p | a _J * (mbar ⁻¹) | \mathbb{R}^2 | p | Average | | | |
| UKRU | 4.26E-02 | 0.83 | < 0.005 | 3.93E-02 | 0.58 | < 0.005 | 2.95E-02 | 0.47 | < 0.05 | 1007.7 | | | |
| UKUB | 1.90E-02 | 0.50 | - | 1.17E-02 | 0.05 | < 0.05 | 4.16E-03 | 0.04 | - | 1011.7 | | | |
| UKRO | 6.33E-02 | 0.95 | < 0.001 | -1.21E-01 | 0.40 | - | -2.98E-02 | 0.17 | - | 1012 | | | |
| GERRU | 5.10E-02 | 0.97 | - | 8.95E-02 | 0.85 | < 0.001 | 2.16E-02 | 0.21 | - | 1007.0 | | | |
| GERUB | 6.27E-02 | 0.97 | - | 4.00E-02 | 0.76 | - | 2.00E-02 | 0.37 | < 0.05 | 995.5 | | | |
| GERRO | 4.57E-02 | 0.79 | - | -9.61E-02 | 0.43 | - | -2.80E-02 | 0.21 | - | 995.5 | | | |
| FINRU | 3.46E-02 | 0.88 | < 0.001 | 2.90E-02 | 0.57 | < 0.001 | 1.05E-02 | 0.14 | - | 985.1 | | | |
| FINUB | 2.61E-02 | 0.55 | < 0.005 | -3.57E-03 | 0.02 | - | 4.38E-03 | 0.05 | - | 1004.4 | | | |
| FINRO | 4.91E-02 | 0.70 | - | -2.67E-02 | 0.17 | - | 1.43E-02 | 0.26 | - | 1008.8 | | | |
| SPARU | -2.02E-02 | 0.09 | - | 4.79E-02 | 0.14 | - | 2.89E-02 | 0.08 | - | 939.3 | | | |
| SPAUB | -2.83E-02 | 0.44 | < 0.05 | 1.86E-02 | 0.08 | - | 1.68E-02 | 0.21 | - | 1006.3 | | | |
| GRERU | 6.00E-02 | 0.46 | < 0.001 | -1.50E-01 | 0.73 | - | 8.14E-02 | 0.33 | - | 1014.5 | | | |
| GREUB | 9.42E-03 | 0.10 | < 0.05 | -1.00E-01 | 0.71 | - | 1.58E-02 | 0.04 | - | 1015.7 | | | |





Table 4: Normalised slopes (non-normalised for growth rate), R^2 and p-values (- for values >0.05) for the relation between atmospheric composition variables and NPF event variables.

| | $\mathrm{SO}_2(\mu\mathrm{g}\;\mathrm{m}^{-3})$ | | | | | | | | | | | | |
|-------|---|----------------|---------|---------------------------|----------------|---------|-----------------------------|----------------|---------|---------|--|--|--|
| Site | $a_N* (\mu g^{-1} m^3)$ | \mathbb{R}^2 | p | \mathbf{a}_{G} | \mathbb{R}^2 | p | $a_{J}* (\mu g^{-1} m^{3})$ | \mathbb{R}^2 | p | Average | | | |
| UKRU | -1.97E-01 | 0.38 | < 0.05 | -6.17E-02 | 0.02 | - | 3.30E-01 | 0.06 | - | 1.64 | | | |
| UKUB | -2.57E-01 | 0.62 | < 0.001 | 1.93E-02 | 0.00 | - | 4.18E-01 | 0.40 | - | 2.04 | | | |
| UKRO | -1.03E-01 | 0.82 | < 0.001 | 6.90E-02 | 0.34 | < 0.01 | 8.43E-02 | 0.77 | < 0.001 | 7.46 | | | |
| DENRU | -9.77E-01 | 0.53 | < 0.05 | 2.84E+00 | 0.37 | - | 4.38E-01 | 0.09 | - | 0.52 | | | |
| DENRO | -4.20E-01 | 0.91 | < 0.001 | 6.42E-01 | 0.54 | < 0.005 | 5.66E-01 | 0.62 | < 0.001 | 0.97 | | | |
| FINRU | -5.66E-01 | 0.05 | - | -1.42E+00 | 0.19 | - | -6.30E-02 | 0.00 | - | 0.09 | | | |
| SPARU | -3.62E-01 | 0.74 | < 0.001 | -1.33E-01 | 0.02 | - | -3.55E-02 | 0.01 | - | 0.95 | | | |
| SPAUB | -2.93E-02 | 0.04 | - | 4.12E-01 | 0.59 | - | 1.07E-01 | 0.29 | - | 1.99 | | | |

| NO _x or NO ₂ (ppb) | | | | | | | | | | | | |
|--|---------------------------------------|----------------|---------|---------------------------|----------------|---------|---------------------------------------|----------------|---------|---------|--|--|
| Site | a _N * (ppb ⁻¹) | \mathbb{R}^2 | p | \mathbf{a}_{G} | \mathbb{R}^2 | p | a _J * (ppb ⁻¹) | \mathbb{R}^2 | p | Average | | |
| UKRU | -4.99E-02 | 0.67 | < 0.005 | 4.52E-02 | 0.58 | < 0.05 | -4.51E-02 | 0.70 | < 0.005 | 11.7 | | |
| UKUB | -8.75E-03 | 0.83 | < 0.001 | -3.97E-04 | 0.00 | ı | -1.09E-02 | 0.43 | < 0.05 | 53.6 | | |
| UKRO | -3.22E-03 | 0.72 | < 0.001 | 1.44E-03 | 0.39 | < 0.05 | 2.19E-03 | 0.66 | < 0.001 | 299 | | |
| DENRU | -9.41E-02 | 0.43 | < 0.005 | -4.89E-03 | 0.00 | < 0.001 | -6.47E-02 | 0.55 | < 0.01 | 5.42 | | |
| DENUB | -4.99E-02 | 0.68 | < 0.001 | 2.85E-02 | 0.26 | - | 8.55E-04 | 0.00 | - | 10.5 | | |
| DENRO | -5.10E-03 | 0.75 | < 0.001 | 1.10E-02 | 0.69 | < 0.001 | 8.33E-03 | 0.88 | < 0.001 | 68.5 | | |
| FINRU | -7.27E-01 | 0.54 | < 0.001 | -2.74E-01 | 0.11 | ı | 1.95E-01 | 0.05 | - | 0.72 | | |
| FINRO | -6.24E-03 | 0.68 | < 0.001 | 1.70E-03 | 0.12 | ı | 3.25E-03 | 0.03 | - | 88.1 | | |
| SPARU* | -1.53E-02 | 0.05 | - | 2.54E-02 | 0.01 | ı | 1.25E-01 | 0.21 | - | 3.26 | | |
| SPAUB* | -2.59E-02 | 0.62 | < 0.005 | 2.23E-02 | 0.70 | < 0.001 | 2.57E-03 | 0.01 | - | 31.4 | | |
| GRERU* | 3.01E-01 | 0.19 | - | -1.40E+00 | 0.75 | < 0.001 | 5.23E-01 | 0.13 | - | 0.52 | | |

^{*} NO₂ measurements





| | $O_3(ppb)$ | | | | | | | | | | | | |
|-------|---------------------------------------|----------------|---------|---------------------------|----------------|---------|---------------------------------------|----------------|---------|---------|--|--|--|
| Site | a _N * (ppb ⁻¹) | \mathbb{R}^2 | p | \mathbf{a}_{G} | \mathbb{R}^2 | p | a _J * (ppb ⁻¹) | \mathbb{R}^2 | p | Average | | | |
| UKRU | 2.27E-02 | 0.88 | < 0.001 | -4.89E-02 | 0.53 | < 0.005 | -3.53E-03 | 0.01 | ı | 54.4 | | | |
| UKUB | 1.37E-02 | 0.87 | < 0.001 | -3.45E-02 | 0.68 | < 0.001 | -5.95E-03 | 0.05 | - | 39.3 | | | |
| UKRO | 7.46E-02 | 0.95 | < 0.001 | -1.06E-02 | 0.09 | - | -2.44E-02 | 0.63 | < 0.005 | 16.2 | | | |
| DENRU | 4.97E-02 | 0.92 | < 0.001 | -1.32E-02 | 0.15 | - | 1.23E-02 | 0.08 | ı | 30.1 | | | |
| DENUB | 5.85E-02 | 0.84 | < 0.001 | -1.69E-02 | 0.58 | - | 2.77E-02 | 0.32 | < 0.05 | 28.2 | | | |
| DENRO | 6.42E-02 | 0.51 | < 0.05 | 1.39E-02 | 0.03 | - | 3.24E-02 | 0.91 | < 0.05 | 31.1 | | | |
| FINRU | 6.76E-02 | 0.77 | < 0.05 | -4.23E-02 | 0.60 | - | 3.92E-02 | 0.37 | < 0.05 | 27.4 | | | |
| FINRO | 2.38E-02 | 0.91 | < 0.001 | 6.11E-03 | 0.24 | - | -1.83E-02 | 0.29 | ı | 37.1 | | | |
| SPARU | 1.57E-02 | 0.02 | - | 4.34E-02 | 0.11 | 1 | 1.31E-02 | 0.31 | ı | 75.9 | | | |
| SPAUB | 7.99E-03 | 0.38 | < 0.05 | -5.83E-03 | 0.30 | - | -1.13E-03 | 0.01 | ı | 54.9 | | | |
| GRERU | 7.55E-03 | 0.04 | - | 3.68E-02 | 0.17 | - | -3.01E-02 | 0.15 | - | 49.5 | | | |

| Particulate Organic Carbon (μg m ⁻³) | | | | | | | | | | |
|--|--------------------------------|----------------|---------|---------------------------|----------------|---------|--|----------------|---------|---------|
| Site | $a_{N}^{*} (\mu g^{-1} m^{3})$ | \mathbb{R}^2 | p | \mathbf{a}_{G} | \mathbb{R}^2 | р | a _J * (μg ⁻¹ m ³) | \mathbb{R}^2 | р | Average |
| UKRU | -3.30E-02 | 0.00 | - | 1.13E+00 | 0.42 | < 0.005 | 2.13E-01 | 0.16 | - | 1.96 |
| UKUB | -2.76E-01 | 0.59 | < 0.005 | 6.63E-01 | 0.58 | < 0.05 | 2.19E-01 | 0.55 | < 0.05 | 3.63 |
| UKRO | -3.78E-01 | 0.89 | < 0.001 | 8.12E-01 | 0.57 | < 0.005 | 4.60E-01 | 0.75 | < 0.001 | 6.24 |
| DENRU | -4.44E-01 | 0.75 | < 0.001 | 2.24E-01 | 0.11 | - | -3.17E-01 | 0.68 | < 0.01 | 1.48 |
| DENRO | -7.80E-02 | 0.11 | - | 1.10E+00 | 0.77 | < 0.005 | 4.02E-01 | 0.81 | < 0.005 | 2.59 |
| GERRU | -1.26E-01 | 0.24 | - | 1.35E-01 | 0.09 | - | 3.14E-02 | 0.03 | - | 2.18 |
| FINRU | 2.27E-02 | 0.00 | - | 3.39E-01 | 0.60 | < 0.005 | -3.46E-01 | 0.16 | - | 1.78 |
| GRERU | -2.08E-01 | 0.11 | - | 7.87E-01 | 0.41 | < 0.05 | 8.94E-01 | 0.11 | - | 1.58 |

| Sulphate (μg m ⁻³) | | | | | | | | | | |
|--------------------------------|-------------------------|----------------|---------|---------------------------|----------------|---------|-----------------------------|----------------|--------|---------|
| Site | $a_N* (\mu g^{-1} m^3)$ | \mathbb{R}^2 | p | $\mathbf{a}_{\mathbf{G}}$ | \mathbb{R}^2 | p | $a_{J}* (\mu g^{-1} m^{3})$ | \mathbb{R}^2 | p | Average |
| UKRU ¹ | -2.62E-01 | 0.57 | < 0.001 | 7.34E-01 | 0.77 | < 0.001 | 7.99E-01 | 0.44 | < 0.05 | 1.97 |
| UKUB ¹ | -3.57E-01 | 0.89 | < 0.001 | 9.28E-01 | 0.44 | < 0.01 | 9.72E-01 | 0.16 | - | 1.58 |
| UKRO ¹ | -6.05E-02 | 0.24 | - | 3.04E-01 | 0.34 | < 0.05 | -6.22E-02 | 0.04 | - | 1.98 |
| DENRU ² | -7.81E-01 | 0.34 | < 0.05 | 1.02E+00 | 0.60 | < 0.05 | -1.03E+00 | 0.63 | < 0.01 | 0.52 |
| DENRO ² | -8.23E-01 | 0.28 | - | 1.99E+00 | 0.22 | - | 2.82E-01 | 0.12 | - | 0.55 |
| GERRU ¹ | -3.37E-02 | 0.00 | - | 5.89E-01 | 0.11 | - | -4.89E-02 | 0.01 | - | 0.92 |
| FINRU ³ | -1.18E+00 | 0.65 | < 0.001 | 2.35E-01 | 0.09 | - | -2.53E-01 | 0.17 | - | 1.02 |





- Measurements in PM₁₀
 Measurements in PM_{2.5}
 Measurements in PM₁

| Condensation Sink (s ⁻¹) | | | | | | | | | | | |
|--------------------------------------|----------------------|----------------|---------|---------------------------|----------------|---------|----------------------|----------------|---------|----------|--|
| Site | a _N * (s) | \mathbb{R}^2 | p | \mathbf{a}_{G} | \mathbb{R}^2 | р | a _J * (s) | R ² | р | Average | |
| UKRU | -2.28E+02 | 0.72 | < 0.001 | 2.64E+02 | 0.60 | < 0.001 | 7.58E+01 | 0.22 | - | 3.38E-03 | |
| UKUB | -1.66E+02 | 0.78 | < 0.001 | 2.49E+02 | 0.41 | < 0.05 | 1.73E+02 | 0.35 | < 0.05 | 7.41E-03 | |
| UKRO | -4.03E+01 | 0.75 | < 0.001 | 2.33E+01 | 0.18 | ı | 8.94E+01 | 0.91 | < 0.001 | 2.12E-02 | |
| DENRU | -4.48E+01 | 0.91 | < 0.001 | 6.90E+01 | 0.49 | < 0.05 | 5.37E+01 | 0.24 | - | 9.46E-03 | |
| DENUB | -3.78E+01 | 0.75 | < 0.001 | 3.58E+01 | 0.25 | ı | 1.55E+01 | 0.56 | < 0.005 | 1.42E-02 | |
| DENRO | -1.06E+01 | 0.73 | < 0.001 | 2.53E+01 | 0.56 | < 0.005 | 2.72E+01 | 0.79 | < 0.001 | 3.10E-02 | |
| GERRU | 1.54E+02 | 0.86 | < 0.001 | 1.33E+02 | 0.56 | < 0.001 | 6.67E+01 | 0.63 | < 0.001 | 7.02E-03 | |
| GERUB | 3.59E+01 | 0.56 | < 0.005 | 3.63E+01 | 0.17 | ı | 4.74E+01 | 0.75 | < 0.001 | 9.11E-03 | |
| GERRO | 3.89E+01 | 0.22 | < 0.05 | -2.21E+01 | 0.03 | < 0.005 | 3.54E+01 | 0.45 | < 0.005 | 1.20E-02 | |
| FINRU | -1.80E+02 | 0.59 | < 0.005 | 4.01E+02 | 0.74 | < 0.001 | 4.98E+01 | 0.10 | - | 2.32E-03 | |
| FINUB | -1.51E+02 | 0.63 | < 0.005 | 8.14E+01 | 0.31 | ı | 2.01E+02 | 0.41 | < 0.05 | 6.34E-03 | |
| FINRO | -6.99E+01 | 0.77 | < 0.001 | -1.56E+01 | 0.05 | ı | 2.42E+02 | 0.83 | < 0.001 | 8.96E-03 | |
| SPARU | -2.15E+02 | 0.65 | < 0.005 | 1.86E+01 | 0.00 | ı | 8.60E+01 | 0.47 | < 0.05 | 5.49E-03 | |
| SPAUB | -1.18E+02 | 0.65 | < 0.005 | 3.74E+01 | 0.38 | < 0.05 | 9.51E+01 | 0.52 | < 0.01 | 1.00E-02 | |
| GRERU | 4.33E+00 | 0.00 | - | 2.86E+02 | 0.70 | < 0.001 | 1.77E+02 | 0.56 | < 0.005 | 4.66E-03 | |
| GREUB | 1.64E+02 | 0.65 | < 0.001 | 9.31E+01 | 0.28 | < 0.05 | 1.73E+02 | 0.83 | < 0.001 | 7.55E-03 | |





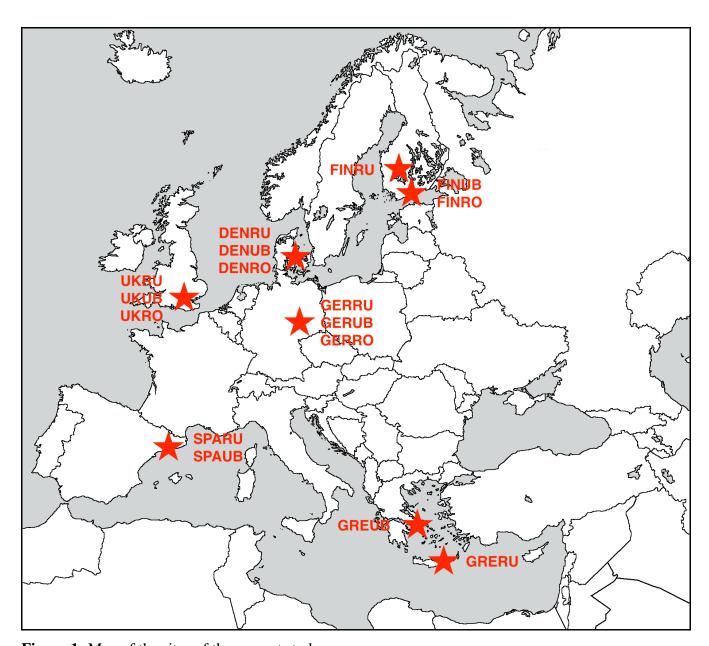


Figure 1: Map of the sites of the present study.





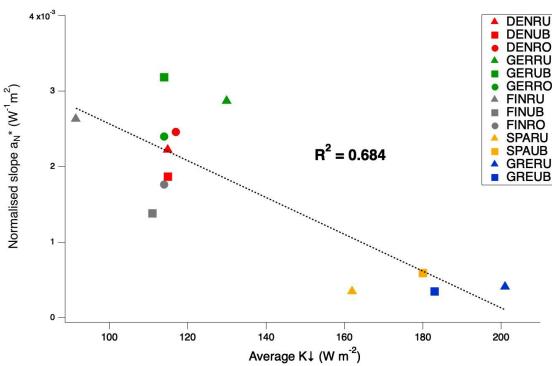


Figure 2: Relation of average downward incoming solar radiation $(K\downarrow)$ and normalised slopes a_N^* for the sites of the present study.

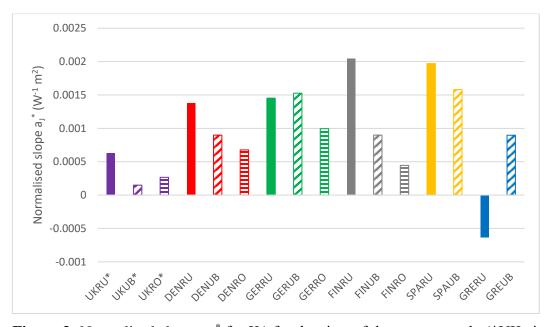


Figure 3: Normalised slopes a_J^* for $K \downarrow$ for the sites of the present study (*UK sites are calculated with solar irradiance).





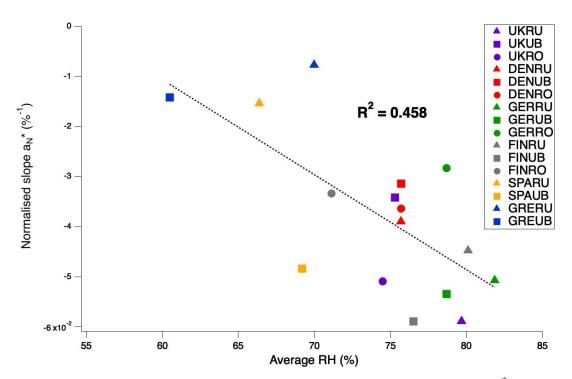


Figure 4: Relation of average relative humidity and normalised slopes ${a_N}^*$ for the sites of the present study.

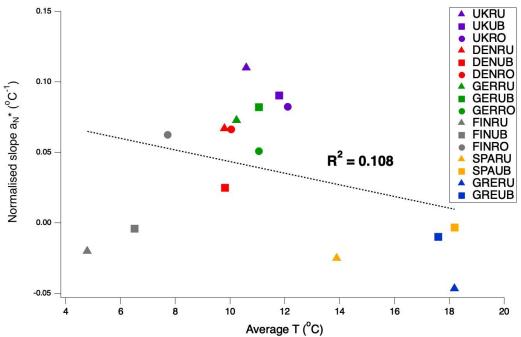


Figure 5: Relation of average temperature and normalised slopes a_N^* for the sites of the present study.





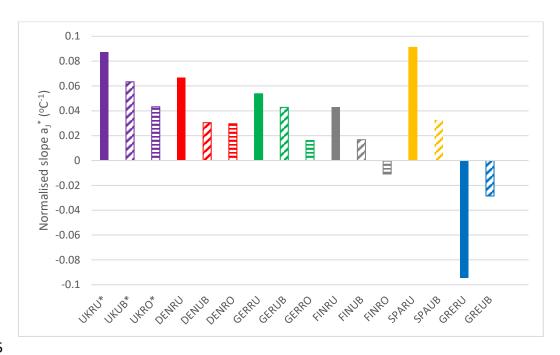


Figure 6: Normalised slopes a_J^* for temperature for the sites of the present study.





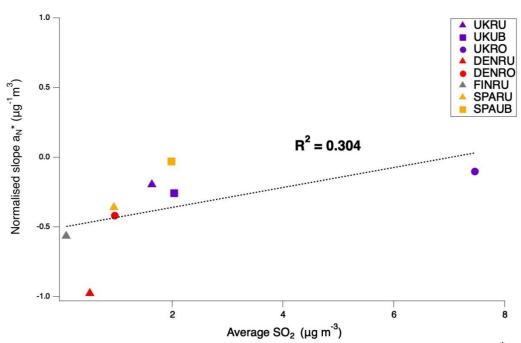


Figure 7a: Relation of average SO_2 concentrations and normalised slopes a_N^* for the sites of the present study.

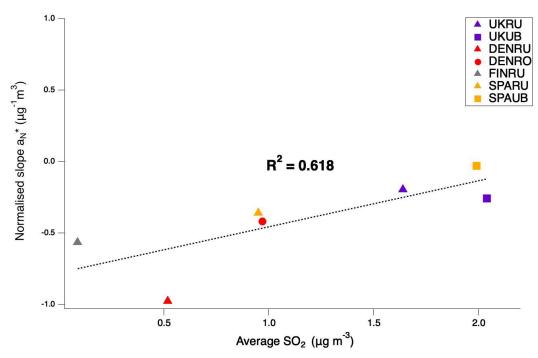


Figure 7b: Relation of average SO_2 concentrations and normalised slopes a_N^* for the sites of the present study (UKRO not included).





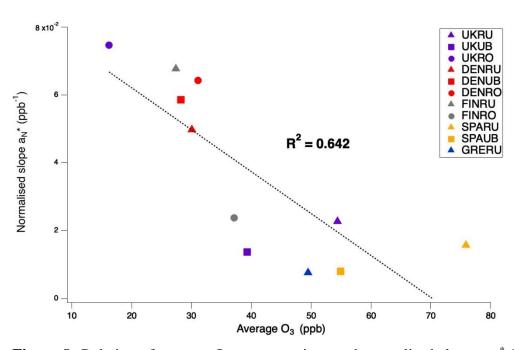


Figure 8: Relation of average O_3 concentrations and normalised slopes $a_N^{\ *}$ for the sites of the present study.

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