

Point by point response to the reviewers:

First of all, we would like to thank the reviewers for the positive and constructive comments. The line number below is indicated based on the clean version. The followings are response to the questions of reviewers.

Comments from the reviewers:

-Reviewer #1

General comments: The manuscript presents an analysis of 24 years of continuous in-situ ground-based methane observations at the remote Mount Waliguan station in China. Such records are very essential for monitoring atmospheric variability and trends of this important greenhouse gas. It is challenging to keep those observations up and running for decades. Thus, I fully support publication and release of these data. However, the analysis of the data is very descriptive and lacks conclusive statements. There is hardly any novel insight into the CH₄ cycle, which is currently drawn from the analysis. Several filters are applied to the data but it is difficult to follow how the filters are applied and which filtered dataset is used in which analysis. The paper may profit from merging chapter 3 (Results) and chapter 4 (Discussion). Moreover, the paper contains quite a number of linguistic flaws. Proofreading by an English native speaker is necessary. Some language mistakes are listed below. The paper is within the scope of 'Atmospheric Chemistry and Physics' but – to my mind – requires major revisions prior to its acceptance for ACP. See my specific comments below.

Response: Thank you very much for your good comments. Based on the comments, our paper has improved a lot. We revised the manuscript and answered the questions point by point.

Overall, i) according to your suggestion, we extended our records from 2017 to 2019, to increase the value and novelty of our study. We updated all our data and updated the respective part in our manuscript including the filtered time series (Table 1, Table 2, and Figure 2), the diurnal variations (Figure 3 and Table S1), the seasonal cycles (Figure 10, Figure 11, and Table S2), and the long-term trends

(Table 4, Figure 12, and Figure S6). Please refer to the revised manuscript for details.

ii) Different data filtering methods were further described and the descriptions of which dataset is used were added. For the meteorological approach, we revised the descriptions as below (Line 197-204, Line 283-285).

“In this study, the CH₄ records associated with local surface winds from selected sectors, i.e. NNE-...-ENE in spring, NE-...-SE in summer, NE-...-ESE in autumn, and NE-ENE in winter, were flagged as locally influenced (27.0%). Subsequently, we rejected portion of daytime records to minimize the effect of human activities (16.9%), e.g. 9:00-13:00 LT (local time) in summer, and 10:00-17:00 LT in winter. Finally, we filtered CH₄ data into locally influenced when the surface wind speeds were less than 1.5 m s⁻¹ to minimize the very local accumulation (9.2%).”

“To precisely understand the characteristics of atmospheric CH₄, including seasonal cycles and long-term trends, it is vital to identify the CH₄ records in well-mixed air without local contaminations.”

For the analysis between the city regions and the Tibetan plateau, the data filtering methods are described as below (Line 205-215).

“In order to investigate the influence of anthropogenic emissions from cities and remote area as Tibetan plateau, we divided the CH₄ data into two main influencing regions according to the analysis, including the geographical conditions, the effect of surface winds, the long-range transports, and the potential source distributions. The first region covers the northeast and southeast (NNE-...SE) of the WLG, which is denoted as City Regions (CR). The second region is located the south to west (S-...-W) of the station and is well known Tibetan (Qinghai-Xizang) Plateau (TP) (Fig. S1). Accordingly, the hourly CH₄ records when the surface winds coming from these sectors were divided into two subsets (i.e. TP and CR). The long-term variations between the two regions as well as the total regional time series were further compared and analyzed.”

For the analysis of the long-range transport of emissions from cities, the data

filtering methods are described as below (Line 665-674).

“As described above, the northeast and southeast city regions might have acted as strong regional sources influencing the atmospheric CH₄ at the WLG. Therefore, to analyze the effect of long-range transport of emissions from cities, the regionally representative data was further excluded by air mass transport, and the remaining regional records were denoted as ‘TR’. First, the monthly cluster analysis was applied to hourly trajectories over 2005-2007, 2008-2012 and 2013-2017. Then, based on the cluster analysis, the clusters were divided into two groups, i.e. from city regions (red clusters in Fig. S8), and other (black clusters in Fig. S8). Finally, the regionally representative data were accordingly classified as two groups based on the cluster results (cities or other). The statistical results were presented in detail in Figure S8 and Table S3.”

For other analysis, i.e. diurnal variation, local surface wind, air mass pathways, and potential source distributions, we used total hourly CH₄ data.

iii) We merged chapter 3 (Results) and chapter 4 (Discussion). The interpretation/conclusion was described in conjunction with results, and more discussions were added. Based on the long-term measurements, we concluded the main findings compared to the existing studies (Line 27-34). Generally, the characteristics of CH₄ varied in different observing periods: i) the diurnal cycle has been becoming apparent and the amplitudes of the diurnal or seasonal cycles increased over time, ii) the wind sectors with elevated CH₄ mole fractions switched from ENE-...-SSE sectors in early periods to NNE-...-E sectors in later years, iii) the area of source regions increased as the years progressed and strong sources shifted from northeast (city regions) to southwest (Northern India), iv) the annual growth rates in recent years (e.g. 2008-2019) were significantly larger than that in early periods (e.g. 1994-2007).

iv) We have asked the language company to help the English improvement.

Specific comments:

Abstract is rather long, you may consider shorten it.

Response: Thank you for your suggestion. We have shortened the abstract (Line

18-34) as below.

“A 26-year long-term record of atmospheric methane (CH₄) measured in-situ at the Mount Waliguan (WLG) station, the only WMO/GAW global station in inland Eurasia, is presented. Overall, a nearly continuous increase of atmospheric CH₄ was observed at the WLG with a yearly growth rate of 5.1 ± 0.1 ppb yr⁻¹ during 1994-2019, except for some particular periods with near-zero or negative values, e.g. 1999-2000, and 2004-2006. The average CH₄ mole fraction was only 1799.0 ± 0.4 ppb in 1994 but increased about 133 ppb and reached a historic level of 1932.0 ± 0.1 ppb in 2019. The case study in the Tibetan Plateau showed that the atmospheric CH₄ increased rapidly. During some special period, it is even larger than that of city regions (e.g. 6.7 ± 0.2 ppb yr⁻¹ in 2003-2007). Generally, the characteristics of CH₄ varied in different observing periods: i) the diurnal cycle has been becoming apparent and the amplitudes of the diurnal or seasonal cycles increased over time, ii) the wind sectors with elevated CH₄ mole fractions switched from ENE-...-SSE sectors in early periods to NNE-...-E sectors in later years, iii) the area of source regions increased as the years progressed and strong sources shifted from northeast (city regions) to southwest (Northern India), iv) the annual growth rates in recent years (e.g. 2008-2019) were significantly larger than that in early periods (e.g. 1994-2007).”

Lines 18-19: write "A 24-year long-term record of atmospheric methane (CH₄) measured in-situ at the Mt. Waliguan station, the only ..., is presented."

Response: Revised the sentence (Line 18-19) as you suggested.

Line 20: "... 1994-2017 ...": why don't you use more recent data (e.g. until end of 2019)?

Response: Thank you for your question. According to your suggestion, we extended our records from 2017 to 2019, to increase the value of our study. We updated all our data and updated the respective part in the manuscript including the filtered time series (Table 1, Table 2, and Figure 2), diurnal variations (Figure 3

and Table S1), seasonal cycles (Figure 10, Figure 11, and Table S2), and long-term trends (Table 4, Figure 12, and Figure S6). Please refer to the revised manuscript for details.

Line 20: "continuously" need to be "continuous"

Response: We have corrected it (Line 20).

Lines 20-22: "continuous increase" but "even negative growth trend ... in particular periods" is a contradiction.

Response: Thank you for the correction. We rewrote it as “a nearly continuous increase” in the revised draft (Line 20).

Lines 24-25: "but unprecedented elevated ~100 ppb" sounds awkward

Response: We updated the data to 2019 and changed it to “but increased about 133 ppb” (Line 24).

Line 24: "... historic high of 1903.8 ± 0.1 ppb in 2016..."; what about 2017?

Response: We have updated it to 2019. “... reached a historic level of 1932.0 ± 0.1 ppb in 2019” (Line 24).

Line 26: what does " $\Delta\text{CO}/\Delta\text{CH}_4$ " mean? Why do you look at CO to CH₄ ratios?

Response: Thank you for your question. ΔCH_4 and ΔCO are the detrended time series of CH₄ and CO based on the method of Thoning et al. (1989). These are results of the original data minus the trend curve. $\Delta\text{CO}/\Delta\text{CH}_4$ is the ratio of the ΔCO and ΔCH_4 . Because part of the CH₄ and CO in atmosphere are from the same anthropogenic sources (e.g. fossil fuel combustion), the long-term trend of $\Delta\text{CO}/\Delta\text{CH}_4$ is helpful to understand the variation of the sources/sinks. We added the descriptions about ΔCH_4 and ΔCO (Line 263-266, Line 464-467) as below.

“The detrended values are denoted as ΔCH_4 and ΔCO , which are the original data points minus the trend curve. To accurately obtain the correlation slopes of

ΔCO and ΔCH_4 , i.e. $\Delta\text{CO}/\Delta\text{CH}_4$, a rolling linear regression was applied to the ΔCH_4 and ΔCO time series by the 'roll_lm' function in 'roll' package of R"

“Because parts of the CH_4 and CO in atmosphere were from the same anthropogenic sources (e.g. fossil fuel combustion), the long-term trend of $\Delta\text{CO}/\Delta\text{CH}_4$ was helpful to understand the variation in the sources/sinks in many studies (Buchholz et al., 2016; Niwa et al., 2014; Tohjima et al., 2014; Wada et al., 2011).”

Line 27: add "elevated" that is reads "... opposite to other elevated sites ..."

Response: We have revised the abstract and deleted the sentence.

Line 30: "(the Northern India)" -> "(Northern India)"

Response: We revised it (Line 32).

Line 35: delete "What is interesting" I do not believe that you want to mention uninteresting things in the abstract of your manuscript.

Response: We deleted it (Line 27).

Line 47: "anomalously" -> "anomalous"

Response: We rewrote the abstract and deleted the related sentence.

Lines 47-49: If this statement is presented prominently in the abstract, it also needs to be discussed in the manuscript. Moreover, when looking at Fig. 11 (lower panel), I do not really see an increasing growth rate (but only a mostly positive growth rate).

Response: Thank you for your comment. Yes, the description is inappropriate here, it should be “larger growth rate” and this is rather a statement that fits to a synthesis analysis. We rewrote the abstract and deleted the related descriptions.

The mostly positive values will lead to larger average growth rate, but the increasing growth rate is not distinct in Fig. 11. We listed the periodic growth rates in Table 4 as below. The growth rates during 2008-2019 are larger than that during

1994-2002.

Table 4. Annual growth rates of atmospheric CH₄ in the City Regions (CR), the Tibetan Plateau (TP), and total regional records during 1994-2019 at the WLG station.

	1994-1997	1998-2002	2003-2007	2008-2012	2013-2019	1994-2019
CR	3.0 ± 0.1	3.6 ± 0.2	5.3 ± 0.2	7.0 ± 0.2	6.2 ± 0.1	5.2 ± 0.1
TP	3.4 ± 0.1	3.0 ± 0.2	6.7 ± 0.2	5.7 ± 0.2	5.7 ± 0.1	5.1 ± 0.1
Total	4.9 ± 0.1	2.5 ± 0.2	4.9 ± 0.1	7.7 ± 0.1	5.5 ± 0.1	5.1 ± 0.1

Line 60: write "CH₄ has an 8-12 years lifetime ..."

Response: We have revised it (Line 45).

Line 64: "... CH₄ rapidly increased ..."

Response: We have revised it (Line 49).

Lines 65-66: write "Results from ice core analyses in Antarctica showed ..."

Response: We revised the sentence (Line 50-51).

Line 67: write "... has reached a level unprecedented over ..."

Response: We have revised the sentence (Line 52).

Line 68: awkward English

Response: We have revised the sentence (Line 53-55) as below.

"At the beginning of the 1990s, the CH₄ mole fraction showed a decreasing trend in. Consequently, the reverse trend has been observed since 1998 due to the higher global mean temperature (Dlugokencky et al., 1998; Nisbet et al., 2014)."

Line 71: delete "special"

Response: We have deleted it.

Line 73: write "... (Nisbet, et al., 2019) followed by a renewed CH₄ increase since then".

Response: We revised the sentence (Line 58).

Line 78: don't start a sentence with "And ..."

Response: We have deleted it.

Lines 78-79: write "unexpected" instead of "not expected"; why the increase is "unexpected"?

Response: We have deleted the related sentence because it is inappropriate here (Line 62). Because the growth rate of CH₄ is originally very low and lasts a long time before 2007, but increases strongly and continuously from 2007 to now. Therefore, we used the 'unexpected increase'.

Line 84: explain "C.E."

Response: We added "Common Era" (Line 66).

Line 85: write "Atmospheric CH₄ is mainly ..."

Response: We have revised it (Line 68).

Lines 105-106: write "Systematic observations are a prerequisite to get an accurate understanding of spatial and temporal behavior of atmospheric CH₄ concentrations."

Response: We revised the sentence (Line 94-95).

Line 113: awkward English

Response: We revised the sentence (Line 101-107) as below.

“Recently, other stations have been installed for CH₄ observation, such as the Barrow (BRW), South Pole (SPO) (polar site), Cape Grim (CGO) and Minamitorishima (MNM) (coastal/island sites), Jungfrauoch (JFJ) and Mount Waliguan (continental mountain site). Hundreds of CH₄ observation stations worldwide are currently running under the framework of the WMO/GAW.”

Lines 114-118: There are many more stations continuously measuring CH₄ levels in the

atmosphere; why did you select these ones?

Response: Thank you for your question. We select these stations because i) they just represent different types of global stations, ii) parts of the stations have similar altitude or latitude to Waliguan, e.g. MLO, JFJ, and MNM.

Lines 118-127: this is largely a repetition, merge with paragraph on the previous page (lines 85-104).

Response: We have moved and merged the sentences (Line 87-93).

Lines 146-147: write "... which is the longest record in China."

Response: We have revised it (Line 126).

Lines 157: write "WLG is the only ..."

Response: We have revised it (Line 136).

Line 158: write "... and is run by the China Meteorological Administration ..."

Response: We have revised it (Line 137).

Line 166: "Tibetan Plateau"

Response: We have revised it (Line 145).

Line 167: "dial variations ... are influenced ..."

Response: We have revised it (Line 146).

Line 221: write "... were flagged as locally influenced."; if the data are locally influenced they are poorly representative.

Response: Thank you for your comments. We have revised it (Line 200). Yes, the data are locally influenced means they are poorly representative. They represent the CH₄ that are strongly affected by local contaminations or not in well-mixed air. We excluded them to get the regionally representative data. Finally, based on the

meteorological method, about 64% of the CH₄ data was classified as regionally representative over 1994-2019 at the WLG (Line 198-204, Table 1).

Lines 220-225: add percentages of rejection according to the individual filters.

Response: We have calculated and added the percentage of each filter (Line 198-204). Overall, about 27.0% by wind sectors, 16.9% from daytime records, and 9.2% by wind speed.

Line 224: "We filtered CH₄ data into local events ..." sounds strange, how can you filter data into events?

Response: We have changed "local events" to "locally influenced". The locally influenced data cannot represent the CH₄ in well-mixed air.

Line 226: write "hourly CH₄ data was binned into 16 horizontal wind direction classes ..."; is this done for all data or the regionally representative data only?

Response: We have revised the sentence (Line 216-217). This is done for all data.

Lines 235-236: elaborate on HYSPLIT: what's the spatial resolution of the model, what is the height above sea level of WLG in the model?

Response: The spatial resolution of the model is 0.5×0.5 degree and the height is 10 km a.s.l. We added the descriptions in the paper (Line 229-230).

"The spatial resolution of the model is 0.5×0.5 degree and the model height is 10 km a.s.l."

Line 239: write "The trajectories for January, April, ..."

Response: We have revised the sentence (Line 231).

Line 285: "appropriately every five years a period" sounds strange.

Response: We have revised the sentence (Line 274-278) as bellow.

"The entire CH₄ time series were divided into five observing periods, i.e. 1994-

1997, 1998-2002, 2003-2007, 2008-2012, and 2013-2019, according to the significant stages or the critical time period of atmospheric CH₄ variations from previous studies.”

Results and Discussion chapters (Chapters 3 and 4): I strongly encourage the authors to merge these two chapters. Results should be discussed in stronger conjunction with existing literature, preferably from Asian sites; I don't see the rationale why some findings are compared with conclusions from non-elevated sites in Europe.

Response: Thank you very much for your constructive suggestions. We have merged the section of 'Results' and 'Discussion'.

We added more discussions with existing long-term studies from Asian sites, e.g. Yonagunijima (YON) and Ryori (RYO) in Japan, Sinhadad (SNG) and Cape Rama station (CRI) over India, Ulaan Uul (UUM) in Mongolia, and Tae-ahn Peninsula (TAP) in Korea. Some added descriptions are as below.

“Tohjima et al. (2014) found an opposite variation at the Hateruma Island, which showed small slope values in the summer. Wada et al. (2011) analyzed more than 10-year seasonal variation of the $\Delta\text{CO}/\Delta\text{CH}_4$ ratios at three monitoring stations, i.e. MNM, Yonagunijima (YON), and Ryori (RYO) in Japan, which also showed an opposite trend to that of the WLG. It was because these sites were considerably affected by the Asian continental source regions, where had enhanced emissions of CH₄ in the summer (rice paddies) and CO in the winter (fuels combustion).” (Line 474-480)

“For other regional sites in the Asia, Guha et al. (2018) studied seasonal variability at the Sinhadad (SNG) and Cape Rama station (CRI) over India, which also showed an opposite trend to the WLG due to the strong impact of monsoon dynamics. Ahmed et al. (2015) found that the seasonal CH₄ showed a maximum in the winter and a minimum in the spring at two urban sites of Guro (GR) and Nowon (NW), in Seoul, Korea over 2004-2013. Kim et al. (2015) investigated the decadal variation (1991-2013) of CH₄ at the East Asian sites, e.g. Ulaan Uul (UUM) in Mongolia and Tae-ahn Peninsula (TAP) in Korea, which revealed again an opposite

seasonal trend to that of the WLG.” (Line 521-529)

“The seasonal amplitude at the WLG (~14 ppb) was significantly lower than many other sites in the Northern Hemisphere, by about 35-70 ppb. Such sites included MLO in America, BRW in North Pole, UUM in Mongolia, TAP in Korea, Ny-Ålesund in Norway, Bialystok in Poland, Ochsenkopf in Germany, and Beromunster in Switzerland (Dlugokencky et al., 1995; Kim et al., 2015; Morimoto et al., 2017; Thompson et al., 2009; Popa et al., 2010; Satar et al., 2016). MBL also showed a larger amplitude than WLG (Fig. 11). The study at the SNG and CRI over India showed a much larger amplitude close to 200 ppb (Guha et al., 2018).” (Line 546-553)

“Tohjima et al. (2002) found that the CH₄ levels at the Cape Ochi-ishi and Hateruma Island in 1995-2000 respectively increased by 4.5 and 4.7 ppb yr⁻¹, which were also similar to that of the WLG. Tsutsumi et al. (2006) analyzed the trend of hourly CH₄ data from 1998 to 2004 on the YON, which showed a similar increase (~3.0 ppb yr⁻¹) to the WLG. The study at the GR and NW in Seoul, Korea, presented almost an identical trend of 2 ppb yr⁻¹ between 2004 and 2013 (Ahmed et al., 2015), which was lower than that of the WLG in similar period.” (Line 586-592)

“The growth rate of CH₄ observed at the Ny-Ålesund, Svalbard, increased from 0.3 ± 0.2 ppb yr⁻¹ during 2000-2005 to 5.5 ± 0.2 ppb yr⁻¹ during 2005-2014, which had a similar variation but with a little lower growth rates than that of the WLG (Morimoto et al., 2017). The study suggested that the temporal pause in 2000-2005 was ascribed to the reductions of CH₄ emissions from the microbial and fossil fuel sectors, while the increase in 2005-2014 was due to an increase in microbial release.” (Line 598-604)

Additionally, we also compared the result with adjacent stations in China, e.g. Shangdianzi, Lin'an. For other sites with similar latitude or altitude in the Northern Hemisphere, e.g. Mauna Loa and Jungfrauoch, are also compared.

Lines 297-298: awkward English

Response: We have revised the sentence (Line 308-309) as bellow.

“In winter, a large increase of CH₄ was found during 9:00-17:00 LT, with the largest peak to trough amplitude of 7.1 ± 2.9 ppb.”

End of Chapter 3.1: interpretation/conclusion missing.

Response: We merged the section of ‘Results’ and ‘Discussion’ and rewrote the Chapter 3.1 (Chapter 3.2 now). The interpretation/conclusion was described in conjunction with results (Line 310-333). Please refer to the revised manuscript.

Line 315: "As observed by the previous short-term observations"; I don't understand this statement.

Response: We have changed to “Similar to the previous studies...” (Line 319).

Line 326: Delete "What interesting is"

Response: We have deleted it.

Line 345: Write "It's obvious that CO showed ..."

Response: We have revised it (Line 391).

Line 348: write "when percentages ranged from 0 to 40."; write "When data exceeded the 60% percentile, the high area probability areas ..."

Response: We have revised the sentence (Line 394-395).

End of Chapter 3.2: interpretation/conclusion missing.

Response: We merged the section of ‘Results’ and ‘Discussion’ and rewrote the Chapter 3.2 (Chapter 3.3 now). The interpretation/conclusion was described in conjunction with results (Line 345-365, Line 374-375, Line 383-391, and Line 395-399). Please refer to the revised manuscript.

Line 365, 368-369: awkward English

Response: We have revised the sentence (Line 414-419) as below.

“Cluster 3 showed the highest CH₄ mole fraction, with an enhancement of ~4 ppb relative to the seasonal average. In winter, the air masses primarily came from northwest and southwest regions, e.g. cluster 3 (59%), and cluster 1 (34%) (Fig. 6d) and the cluster 1 brought the highest CH₄ mole fractions with an enhancement of ~7 ppb over the seasonal average.”

End of Chapter 3.3: interpretation/conclusion missing.

Response: We merged the section of ‘Results’ and ‘Discussion’, and rewrote the Chapter 3.3. The interpretation/conclusion was described in conjunction with results (Line 420-433, Line 451-461).

Line 389: write "regionally representative"

Response: We have revised all the descriptions in the manuscript.

Line 390: write ""locally influenced"

Response: We have revised all the descriptions in the manuscript.

Line 391: "data was ... larger than ... events"; awkward English

Response: We have revised the sentence (Line 288-289) as below.

“The average of the locally influenced data (1868.2 ± 0.3 ppb) was larger than that of the regionally representative records (Table 1).”

Line 394: Does it hold true for all data or regionally representative data?

Response: It hold true for both of them. The average of regionally representative data and all data both showed increasing trend over 1994-2019.

Line 396 (reference to Fig. 7): I suggest to move Fig. 7 above and make it to Fig. 2; show first the whole dataset before you analysis of the data

Response: Thank you very much for your suggestion. We changed Fig. 7 to Fig. 2 and moved Chapter 3.4 above as Chapter 3.1.

Line 402: what is ΔCO and ΔCH_4 ? I assume it is excess CO and excess CH_4 , i.e. data above baseline (looks like when looking at Fig. S1); if so, how was the background determined? Please elaborate; show also the CO time series.

Response: Thank you for your questions. Similar to the question you posed on L26, ΔCH_4 and ΔCO are the detrended time series of CH_4 and CO from 2004-2017 based on the method of Thoning et al. (1989).

The detrended value is the original data points minus the trend curve. The trend value is the polynomial part of the function plus the long-term filter of the residuals. We have added the descriptions in the manuscript (Line 261-266).

The ΔCH_4 and ΔCO time series are showed in Figure S4. The original hourly CO time series was also added in the paper (Figure S3).

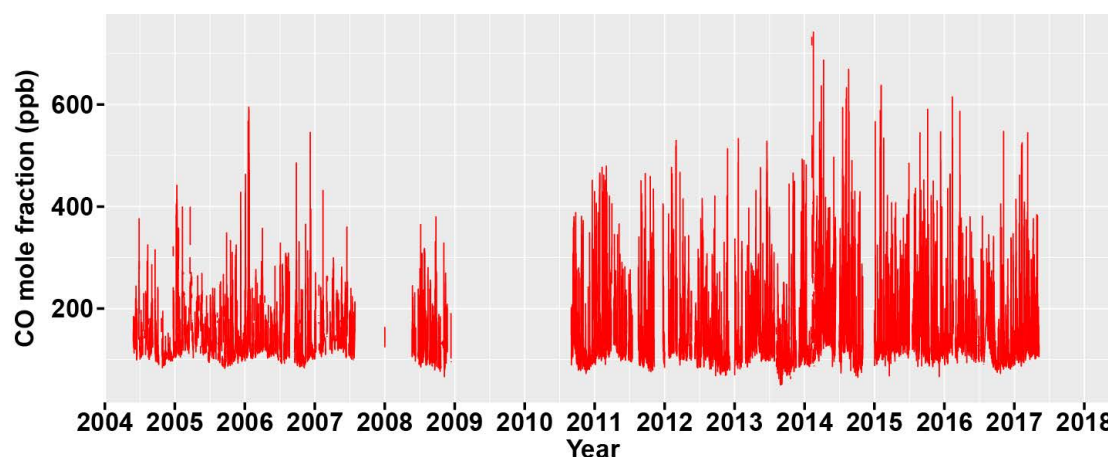


Figure. S3. The hourly CO data from 2004 to 2017 at the WLG station.

Line 411: incomplete sentence.

Response: We deleted the sentence because there is a repetition of above descriptions.

End of Chapter 3.5: what do we learn from the different ΔCO to ΔCH_4 ratios?

Response: The strong seasonal variation of $\Delta\text{CO}/\Delta\text{CH}_4$ also revealed that the WLG was affected by different anthropogenic sources, e.g. sources from cities, and sources from the Tibetan Plateau, during a year, especially in the summer and

winter. The long-term trend of the slopes implied that the source emission types (CO sources or CH₄ sources) around the WLG might have been changing with human activities, like straw burning, in the early years, or coal mining in recent years. We revised the Chapter 3.5 and added the descriptions (Line 498-504). Please refer to the revised manuscript.

Lines 415-419: this is a different filtering than the one described above, right? Move these lines to Chapter 2.3.

Response: Yes, this is a different data filtering method. We are using this method to further investigate the influence of anthropogenic emissions from cities and remote area as Tibetan plateau. We moved the paragraph to Chapter 2.3 (Line 205-215).

Line 426: write "... the seasonal averages of regionally representative CH₄ were ..."; do these numbers relate to regionally representative data?

Response: We have revised it. Yes, the seasonal averages are calculated based on regionally representative data.

End of Chapter 3.6.1: interpretation/conclusion missing.

Response: We merged the section of 'Results' and 'Discussion' and rewrote the Chapter 3.6.1. The interpretation/conclusion was described in conjunction with results (see Chapter 3.6.1, Line 518-562).

Line 442: write "... growth rates were very small or even negative ..."

Response: We have revised it (Line 567).

Lines 466-468: How is this filtering exactly done? Please explain. How many data (in %) are remaining?

Response: The regionally representative data are further filtered based on the hourly trajectories of cluster analysis.

First, the monthly cluster analysis was applied to hourly trajectories over 2005-2007, 2008-2012 and 2013-2017. Then, based on the cluster analysis, the clusters were divided into two groups, i.e. from city regions (red clusters in Fig. S8), and other (black clusters in Fig. S8). Finally, the regionally representative data were accordingly classified as two groups based on the cluster results (cities or other). We added the descriptions in the manuscript (Line 669-674).

Figure S8, Table S3, and Table 5 showed the filtered results in detail. Eventually, 67.1% of regionally representative data are remaining.

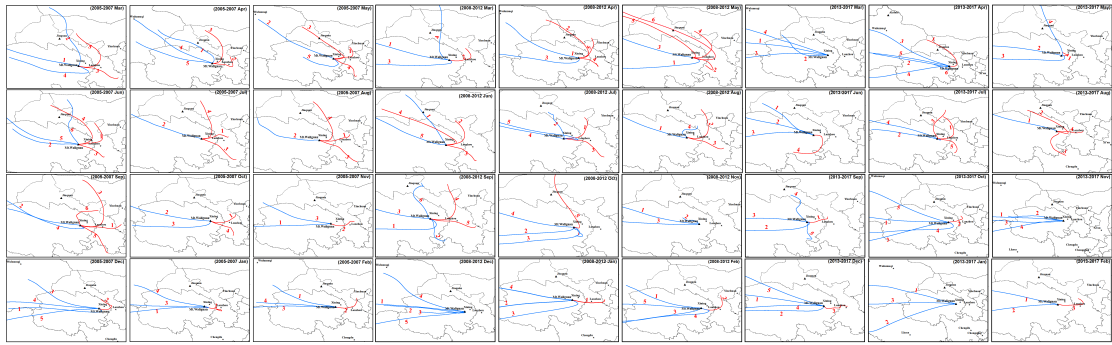


Figure S8. The monthly cluster analysis of hourly trajectories during 2005-2017 at the WLG station. The red clusters represent air masses transport from city regions.

Table S3. The statistics of the excluding trajectories over different observing periods during 2004-2017 at the WLG station.

		Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Oct	Nov	Dec	Sum
2005-2007	Total	764	1193	938	1190	882	1139	921	1431	1427	843	1272	911	12911
	Reject	157	111	51	371	545	767	632	1325	911	181	133	19	5203
2008-2012	Total	3060	2081	1764	1811	1136	2250	1919	2826	2791	1527	1868	1079	24112
	Reject	41	228	194	467	675	1681	1428	1727	1298	90	0	0	7829
2013-2017	Total	2110	1738	2112	1882	1269	1259	461	1816	2186	1659	1844	1111	19447
	Reject	0	1	0	55	81	0	90	875	138	71	0	12	1323

Table 5. The statistics of CH₄ data without air mass transport from city regions (TR) over different periods during 2005-2016 at WLG station.

	Transport regions	Hours	Percentage (%)	Average (ppb)	Updated growth rate (ppb yr ⁻¹)
2005-2007	TR	6922	77.2	1824.9 ± 0.2	2.7 ± 0.2
	City	2041	22.8	1835.9 ± 0.5	-

2008-2012	TR	7060	64.4	1853.7 ± 0.2	10.1 ± 0.1
	City	4254	35.6	1861.0 ± 0.3	-
2013-2016	TR	4152	61.6	1888.2 ± 0.3	6.3 ± 0.1
	City	2591	38.4	1888.5 ± 0.5	-
2005-2016	TR	18134	67.1	1850.6 ± 0.2	7.0 ± 0.1
	City	8886	32.9	1863.2 ± 0.3	-

Line 481: "no significant trend was found ..." what does that mean? How do you interpret this finding?

Response: It means that the CH₄ from the Tibetan Plateau has similar growth rates to that from city regions. We revised the description (Line 581-582) as below.

“Similar growth rates were found between the CR and the TP during 1994-2019 (Fig. S6).”

The finding implied that i) the atmospheric CH₄ at the WLG was not predominantly influenced by eastern cities in recent years and ii) large amounts of CH₄ were transported from the Tibetan Plateau to WLG in recent years (Line 699-702). Through the analysis of spatial distribution of emissions (Fig. 7, Fig. S10), the emission from southwest to south regions, i.e. Northern India, increased strongly after 2007, which may have contributed a lot to the high growth rates in the Tibetan Plateau in recent years.

Line 495: Write "WLG was increasingly affected by local sources ..." which ones are these?

Response: We revised the sentence (Line 331-333).

“The increasing amplitude and CH₄ mole fractions suggested that the WLG was increasingly affected by local and regional anthropogenic sources, such as gas exploitation, and grazing.”

Line 500: "... higher CH₄ ... in past years"; which years, be more explicit.

Response: We revised the descriptions, “In 1994-2002, the CH₄ mole fractions ...” (Line 310).

Line 505: this statement is trivial as Zhou et al. used the same data, right?

Response: Zhou et al. (2004) used the data from 1991 to 2002. However, we are using much longer observation records than them. The compared period is also not absolutely the same. Through the 26-year measurements, we found different results, which provided a broader view of changing characteristics of CH₄ to understand CH₄ cycle.

In this chapter 3.3, compared to Zhou et al. (2004), we found that wind sectors with elevated CH₄ mole fractions changed over time (Line 366-375). The elevated CH₄ was predominately from the ENE-...-SSE sectors in the early years (Fig. 4a-b) but evolved to the NNE-NE-ENE-E sectors in later years (Fig. 4c-e).

Line 509: "... which could emit large amount of CH₄ by human activities"; write "large amounts"; statement is very vague, can you be more specific?

Response: We have revised the sentence as below (Line 353-359).

“The two largest cities of Xining (with population of ~2.2 million) and Lanzhou (with population of ~4 million) are also situated in the northeast and east of the WLG, respectively. The heavy human activities from anthropogenic fossil combustion, landfills, and livestock, could have also emitted large amounts of CH₄. Based on the data from the Emissions Database for Global Atmospheric Research (EDGAR), the increase of CH₄ emission was 500 kg yr⁻¹ in these two regions throughout 1994-2015 (Crippa et al., 2019).”

Line 512: write "... may also contribute to the high CH₄ ..."; statement is very vague, can you be more specific?

Response: We have revised the sentence as below (Line 359-365).

“Also, the Yellow River Canyon industrial area (YRC), which is ~500 km northeast of the WLG, may also have contributed to the high CH₄ values (Zhou et al., 2003). With the rapid development of land-use, water utilization, and agriculture sources in the YRC, large CH₄ emissions could have easily transported to the WLG. Previous studies on black carbon (BC) and carbon monoxide (CO)

also revealed that the high CH₄ values at the WLG in winter were a result of transport from the YRC (Tang et al., 1999; Zhou et al., 2003).”

Chapter 4.1 misses quantitative statements; what’s new compared to existing literature?

Response: We merged Chapter 4.1 with the ‘Results’ and added more descriptions (see Chapter 3.2 and 3.3). The interpretation/conclusion was added in conjunction with results.

Compared to previous studies, we found that i) the diurnal variations were ambiguous before 2002, but significant diurnal cycles appeared afterward, ii) the highest CH₄ mole fraction was found in winter over 1994-2002, but in summer during the periods of 2003-2007, 2008-2012, and 2013-2019, iii) the wind sectors with high CH₄ mole fractions changed and concentrated on ESE-ENE sectors, and iv) the amplitude of enhancements was increasing.

Line 530: "It was possibly due to ..."; very vague statement, can you be more conclusive?

Response: We have revised the statement (Line 424-433).

“The higher values seen in the northwest to southwest airflow were because the air masses had passed through the northwest of the Qinghai province and the central area of the Xinjiang Uygur Autonomous Region (XUAR). This is where the Ge’ermu urban area (the second largest city of Qinghai) was located and where there was with rapid industrial development, natural gas and petroleum resource exploitation, and residue burning of large crops, hence the CH₄ emissions were strong (Fang et al., 2013; Zhang et al., 2013). This result was also similar to the CPF percentile analysis (Fig. 5), as time went by, the southwest or northwest regions that were farther away from the site became the strongest source regions to the WLG.”

Line 536: "the southwest or northwest region ... may be also strong source regions"; very vague statement.

Response: We have revised it as “This result was also similar to the CPF percentile analysis (Fig. 5), as time went by, the southwest or northwest regions that were farther away from the site became the strongest source regions to the WLG.” (Line 431-433).

Line 540-542: awkward English; no conclusive statement.

Response: We have revised the sentence (Line 451-453) as below.

“More CH₄ sources appeared at the WLG along with the progression of time, which could have been attributed to the influence of human expansion.”

Lines 544-545: add reference.

Response: We added the reference (Fu et al, 2012) (Line 457).

Line 556: "It is of great possible ..." awkward English

Response: We have deleted it.

Line 573: Hateruma Island is at sea level, why do you compare WLG with this station?

Response: We compared WLG with Hateruma Island (Tohjima et al, 2014), because the studies they reported about the long-term slope of $\Delta\text{CO}/\Delta\text{CH}_4$, and we used similar method by Tohjima et al. (2014) to calculate $\Delta\text{CO}/\Delta\text{CH}_4$. In addition, as an island station, it could well capture the CH₄ signal that is unaffected by human activities.

Line 582: awkward English

Response: We have merged Chapter 4.2 with ‘Results’, the statement is a repetition and we deleted it.

Chapter 4.2 is very descriptive and lacks conclusions

Response: We merged the section of ‘Results’ and ‘Discussion’ and added interpretation/conclusion in conjunction with result (see Chapter 3.4.1 and 3.4.2).

Lines 592-597: move in front of line 591.

Response: We have moved the sentence (Line 530-535).

Line 602: start station names with upper case letters; why do compare with these stations, which have different characteristics.

Response: We revised them (Line 546-550). The seasonal amplitude at WLG was lower than many other sites in the Northern Hemisphere. We compared these stations because they are typical stations with large seasonal amplitude. We also added more discussions to compare the study with Asian sites.

“The seasonal amplitude at the WLG (~14 ppb) was significantly lower than many other sites in the Northern Hemisphere, by about 35-70 ppb. Such sites included MLO in America, BRW in North Pole, UUM in Mongolia, TAP in Korea, Ny-Ålesund in Norway, Bialystok in Poland, Ochsenkopf in Germany, and Beromunster in Switzerland (Dlugokencky et al., 1995; Kim et al., 2015; Morimoto et al., 2017; Thompson et al., 2009; Popa et al., 2010; Satar et al., 2016). MBL also showed a larger amplitude than WLG (Fig. 11). The study at the SNG and CRI over India showed a much larger amplitude close to 200 ppb (Guha et al., 2018).”

In order to fully discuss the characteristic of CH₄ at WLG, based on the existing studies, we compared WLG with both adjacent stations around China and other stations with similar altitude or latitude in the Northern Hemisphere, e.g. Asian sites, European sites, and American site.

Line 611: why is the photochemical capacity weak? I would expect a high photochemical activity.

Response: Yes, the photochemical capacity is high in summer. We have revised the statement (Line 539-542). But as a plateau station with lower VOC and O₃, the oxidizing capacity at WLG is far below the lower altitude area.

Lines 630-631: numbers were already given in lines 454 ff.; repetition won't be needed

if Results and Discussions are merged.

Response: Thank you for your good suggestion. We have merged 'Results' and 'Discussion'.

Lines 644-645: why do you compare with a Swiss site? You may compare with composite numbers, e.g. from the WDCGG data summary report (most recent version is #43); downloadable on the WDCGG webpage.

Response: Thank you for your comments. We have compared the data from WDCGG sum43 as well as WMO GHG Bulletin, e.g. (Nisbet et al., 2016; 2019), (WMO, 2019; 2020) (Line 605-613). The comparisons of other sites with similar altitude or latitude are also listed, aimed to fully discuss the long-term CH₄ growth in the world. We also added more discussions about the Asian sites (Line 586-592 and 598-604).

Lines 656-657: are the conclusions drawn for the Swiss site also true for WLG?

Response: To our knowledge, until now, the reasons of the anomalous spikes or strong growth of atmospheric CH₄ have not yet been determined. The conclusion raised by the study at Beromünster, Swiss is also a possible reason. However, we are trying to provide more information to figure the cause for these spikes.

We analyzed the long-term variation between city regions and Tibetan Plateau, the CH₄ emission from different sectors in China, and source changes over different periods. We also discussed the potential reasons provided by other studies (Line 631-645). We concluded that i) the emission from solid fuel (e.g. coal) and rice cultivation may contribute to the anomalous increase (Line 649-653), ii) large emissions from Northern India in recent years may have contributed to the anomalous increase at the WLG (455-461), iii) the warming in the Tibetan Plateau is also an important factor (Line 693-695).

Lines 658-680: this paragraph may better fit into the introduction.

Response: Thank you for your comments. The paragraph describes the CH₄

emissions based on the data from EDGAR. It could fit into the introduction. But based on previous studies, we already generally described CH₄ sources using one paragraph in the introduction (69-93).

This paragraph particularly describes the CH₄ emissions from different sectors during 1995-2015 in China, which showed a similar period to our study. The data could be used to discuss the potential reasons for CH₄ growth at the WLG. Hence, we merged the paragraph with 'Results' (Chapter 3.6.2).

Lines 691-692: "suggested that there were possibly other strong CH₄ sources ...": which ones? The following lines do not provide any answer.

Response: Based on our study, it's the strong sources from Northern India. India with abundant cattle as well as an extensive large-scale coal mining operation possibly contributed large amounts of CH₄ to move from northern India to the northeastern Tibetan Plateau (Fig. 7i & l). We have revised the sentence as below (Line 685-687).

"These results suggested that there were possibly other strong CH₄ sources at the WLG that were not from cities and the southwest region (Northern India) was the most likely contributor."

Line 702: "due to the emission from the two largest source regions": be more specific.

Response: We have revised the sentence as below (Line 697-699).

"This would be especially true with the scenario of quickly increasing CH₄ on the Qinghai-Tibetan Plateau due to the emissions from the two largest source regions of Northern India and Eastern China."

Lines 704-705: I don't understand what wants to be said here.

Response: We would like to show the importance of studying carbon cycle in the Tibetan Plateau. We merged the 'Results' and 'Discussion'. The sentence is not appropriate here. Hence, we deleted it and added the descriptions as below (Line 696-702).

“The rapid increase of CH₄ would probably make it difficult to meet the goals of carbon emission reduction in the future. This would be especially true with the scenario of quickly increasing CH₄ on the Qinghai-Tibetan Plateau due to the emissions from the two largest source regions of Northern India and Eastern China. The large growth rate of atmospheric CH₄ in the TP revealed that i) the atmospheric CH₄ at the WLG was not predominantly influenced by eastern cities in recent years and ii) large amounts of CH₄ were transported from the Tibetan Plateau to WLG in recent years.”

Line 705: "anomalously" -> "anomalous"

Response: We have revised it.

Lines 706-707: how about trends in other regions in the world? This statement is the same as in the abstract and I still don't understand it. This is rather a statement that fits to a synthesis analysis as it is e.g. done in the annual WMO GHG bulletin.

Response: The trend at WLG is similar to other stations in the Northern Hemisphere, and also the global level reported by WMO and WDCGG sum43, especially after 2007, with high growth rate.

Thank you for your comment. Yes, the statements are not appropriate here, and they fit to a synthesis analysis better. We rewrote the abstract and conclusion and revised the related sentence (Line 715-719).

“Additionally, the Tibetan Plateau was intensively affected by strong sources over time, which showed a larger growth rate than that of the city regions in some periods. The anomalous variation and unprecedented growth rate of the atmospheric CH₄ in this region revealed that it was urgent to control CH₄ emissions.”

Lines 722-723: "... the long-term verification is extremely important to ... understand CH₄ variations ..."; did the understanding improve based on the present analysis? Which are the lessons-learnt?

Response: We have revised the conclusion and deleted the related sentence.

The 26-year measurements provided a broader view of changing characteristics of CH₄ at WLG and improved our understanding of the future trend, such as three developing stages of CH₄ (Fig. 12a), rather than a limited view on CH₄ growth, e.g. steady or negative growth in 1994-2006. We also discussed the potential reasons for the increasingly long-term trend (Chapter 3.6.2 and 3.7).

In this study, we found new characteristics of CH₄ at WLG. Generally, the characteristics of CH₄ varied in different observing periods: i) the diurnal cycle has been becoming apparent and the amplitudes of the diurnal or seasonal cycles increased over time, ii) the wind sectors with elevated CH₄ mole fractions switched from ENE-...-SSE sectors in early periods to NNE-...-E sectors in later years, iii) the area of source regions increased as the years progressed and strong sources shifted from northeast (city regions) to southwest (Northern India), iv) the annual growth rates in recent years (e.g. 2008-2019) were significantly larger than that in early periods (e.g. 1994-2007).

Lines 725-726: "Tibetan Plateau was with the highest average altitude ..." awkward English.

Response: We have deleted the description.

Line 727: "anomalously" -> "anomalous"

Response: We have revised it (Line 717).

Data availability statement: very strangely, this paragraph only refers to the data from the other stations but nothing is said how to access WLG data. The data used in the analysis doesn't seem to be freely available since WDCGG only contains daily and monthly CH₄ averages from WLG; where can anybody access hourly CH₄ data from WLG?

Response: Thank you for your question. We will upload the data as supplementary material.

-Reviewer 2

General comments: In this manuscript, a 24-year long-term observation of atmospheric CH₄ at Waliguan WMO/GAW global station in the Tibetan Plateau was studied. This report is very meaningful and the analysis of the paper is very comprehensive. The CH₄ variations and its related potential causes during the long-term observation have been analyzed in details, which would help the scientific community to understand carbon cycle and formulate more informed carbon reduction policy. Given the importance and value of the long-term measurements of CH₄ in the Tibetan Plateau plus that this manuscript is well drafted, I would recommend accepting this paper after minor modifications listed as below.

Response: Thank you very much for your positive comments. We revised the manuscript and answered the questions point by point.

Specific comments:

Line 98: add 'About' before 90%.

Response: We have added it (Line 82).

Line 233-238: How frequent is the backward trajectory computed? Hourly? Please specify the numbers of trajectories are determined.

Response: Yes, it's hourly. We computed the 3-day back trajectories coincident with hourly CH₄ mole fractions. The number of trajectories was determined by the valid hourly CH₄ observations. We added the descriptions (Line 229-230). The related information, e.g. numbers, and average mole fractions, was listed in Table 3.

Some expression is not professional, e.g. "long-distance transport" (Lines 465), can be expressed by "long-range transport", and so on. Please polish the whole report.

Response: Thank you for your suggestion. We have revised all the descriptions in the manuscript.

There are several places in results and discussion where too many details are given, which make the text a little difficult to follow. Results and discussion are suggested to be merged.

Response: Thank you very much for your constructive suggestion. We merged 'Results' and 'Discussion' and rewrote the Chapter. The interpretation/conclusion was described in conjunction with results.

Lines 685-686: Suggest to give more discussions about the larger growth rates in the Tibetan Plateau than that of city region.

Response: We rewrote the Chapter 3.7. The interpretation/conclusion was described in conjunction with results. More discussions were added (Line 685-702).

“These results suggested that there were possibly other strong CH₄ sources at the WLG that were not from cities and the southwest region (Northern India) was the most likely contributor. The PSCF analysis also supported this result (Fig. 7). At present, Northern India and Eastern China were the two largest sources of CH₄ at the WLG (Fig. S10) (Crippa et al., 2019). Since the Tibetan Plateau was coincidentally trapped in the middle of them, the atmospheric CH₄ at WLG was very likely dominated by long-range transport from these two regions. Although CH₄ emissions increased slowly during 1994-2002, a negative trend appeared (Fig. S10), significantly increased emissions were found in both southeast and southwest Asia after 2007. Chen et al. (2013) illustrated that the warming (0.2 °C per decade) in the Tibetan Plateau resulted in substantial emissions of CH₄ due to the thawed permafrost and melted glaciers. The rapid increase of CH₄ would probably make it difficult to meet the goals of carbon emission reduction in the future. This would be especially true with the scenario of quickly increasing CH₄ on the Qinghai-Tibetan Plateau due to the emissions from the two largest source regions of Northern India and Eastern China. The large growth rate of atmospheric CH₄ in the TP revealed that i) the atmospheric CH₄ at the WLG was not predominantly influenced by eastern cities in recent years and ii) large amounts of CH₄ were transported from the Tibetan Plateau to WLG in recent years.”

Line 1057: delete 'CO'.

Response: We are sorry for the mistake. The average CO mole fractions were added in Table 3.

Why are the points not on the line in figure 8 and 9?

Response: Thank you for your question. The lines are the smoothed curves of the points using the method of 'LOESS' Curve Fitting (Local Polynomial Regression), hence they do not just connect the points.

1 **Changing characteristics of atmospheric CH₄ ~~i~~on the Tibetan Plateau~~;~~,**
2 **records from 1994 to ~~2017~~2019 at the Mount Waliguan station**

3

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18 **Abstract.** A ~~264~~-year long-term ~~observation record~~ of atmospheric methane (CH₄)
19 ~~was presented~~measured in-situ ~~at the~~ Mount ~~t.~~Waliguan (WLG) station, the only
20 WMO/GAW global station in inland ~~of~~ Eurasia, is presented. Overall, ~~a~~ ~~during 1994-~~
21 ~~2017,~~nearly continuously increase of atmospheric CH₄ was observed at the WLG with
22 a yearly growth rate of 5.1 ± 0.1 ppb yr⁻¹; during 1994-2019, ~~although near-zero and~~
23 ~~even negative growth appeared in~~ except for some particular periods with near-zero or
24 negative values, e.g., e.g. 1999-2000, and 2004-2006. The average CH₄ mole fraction
25 was only ~~1799~~805.08 ± 0.41 ppb in 1994~~5~~, but ~~unprecedented elevated~~ increased
26 about 100-133 ppb and reached a historic high level of ~~1903~~2.08 ± 0.1 ppb in ~~2016~~2019.
27 ~~The seasonal averages of atmospheric CH₄ at WLG were ordered by summer, winter,~~
28 ~~autumn and spring, and the correlation slopes of $\Delta\text{CO}/\Delta\text{CH}_4$ showed a maximum in~~
29 ~~summer and minimum in winter, which was almost opposite to other sites in the~~
30 ~~northern hemisphere, e.g., Mauna Loa, Jungfrauoch, and was caused by regional~~
31 ~~transport. The case study in the Tibetan Plateau showed that the atmospheric CH₄~~
32 ~~increased rapidly. During some special period, it is even larger than that of city regions~~
33 ~~(e.g. 6.7 ± 0.2 ppb yr⁻¹ in 2003-2007). Generally, the characteristics of CH₄ varied in~~
34 ~~different observing periods~~Strong potential sources at WLG were predominately
35 ~~identified in northeast (cities, e.g., Xining, Lanzhou) and southwest (the Northern~~
36 ~~India), and air masses from west and northwest regions were accompanied with higher~~
37 ~~CH₄ mole fractions than that from city regions.~~

38
39 ~~What is interesting is that obviously changes appeared in different observing periods:-~~
40 ~~Generally,~~ i) the diurnal cycle has been becoming apparent and the amplitudes of the
41 diurnal or seasonal cycles increased over time~~were continuously increasing over time~~,
42 ii) the wind sectors with elevated CH₄ mole fractions switched~~moved~~ from ENE-...-
43 SSE sectors in early periods to NNE-...-E sectors ~~(city regions)~~ in later years, iii) the
44 area of source regions ~~was increased~~edging along with the years, progressed and strong
45 sources ~~gradually~~ shifted from northeast (city regions) ~~to~~ southwest (Northern India),

46 iv) the annual growth rates in recent years (~~e.g., e.g. 2008~~~~13-2016~~2019) were
47 significantly ~~significantly~~ larger than that in early periods (~~e.g., e.g. 1998~~~~1994-~~
48 ~~2012~~2007). ~~We conclude that the site was more and more affected by regional sources~~
49 ~~along with the time. Northern India was possibly becoming the strongest source area to~~
50 ~~WLG rather than city regions before. The case study in the Tibetan Plateau showed that~~
51 ~~the atmospheric CH₄ observed in Qinghai Tibetan Plateau changed not as expected, the~~
52 ~~annual growth rate was even larger than that in city regions in some period (e.g., $7.3 \pm$~~
53 ~~0.1 ppb yr^{-1} in 2013-2016). It is unambiguous that the anomalously fluctuations of~~
54 ~~atmospheric CH₄ in this region are a warning to the world, its increasingly annual~~
55 ~~growth rate may be a dangerous signal to global climate change.~~

56 1 Introduction

57 Since the pre-industrial era, the emissions of greenhouse gases (GHGs) have increased
58 continuously, and larger absolute increases were ~~have been~~ found in recent years, with
59 the concentration now being higher than ever ~~now~~ (WMO, 2019). The GHGs ~~could~~
60 perturb the infrared radiation balance, which traps the heat in the atmosphere and,
61 ~~which~~ contributes to global warming, melting glaciers, extreme weather events and
62 many other global climate changes (IPCC, 2014). The recent 30-year span from 1983
63 to 2012 ~~was~~ the warmest of the last 800 years in the Northern Hemisphere, and half
64 of the rising surface temperatures ~~were~~ due to increased GHGs emissions (IPCC,
65 2014). As one of the most important GHGs, methane (CH₄) has ~~the~~ global warming
66 effect ~~that of methane (CH₄)~~ is just less than after carbon dioxide (CO₂) (Etminan et al.,
67 2016). CH₄ has an 8-12 years ~~atmospheric~~ lifetime (Battle et al., 1996), with ~~the a~~
68 global warming potential ~~of~~ ~23 times greater than CO₂ over a 100-year horizon
69 (Weber et al., 2019). CH₄ ~~about~~ also contributed to about 17% of the radiative forcing
70 caused by long-lived greenhouse gases ~~was contributed by from CH₄ during~~ 1750 to -
71 2016 (Etminan et al., 2016). Since the beginning of the industrial era, the concentration
72 of CH₄ ~~is~~ rapidly increased because of the influence of anthropogenic activities
73 (Saunois et al., 2016). ~~The~~ Results from by the analyses of ice cores analyses in
74 Antarctica showed that the atmospheric concentration of CH₄ ~~has~~ reached a level
75 unprecedented ~~unprecedented~~ over the last 0.8 million years (IPCC, 2014).

76 At the beginning of the 1990s, the CH₄ mole fraction showed a decreasing trend.
77 Consequently, the reverse trend has been observed since 1998 due to the higher global
78 mean temperature ~~In the beginning of 1990s, CH₄ concentration appeared a decreasing~~
79 ~~trend in global scale. However, high growth rates were found in 1998, which was~~
80 ~~possibly due to the higher global mean temperature~~ (Dlugokencky et al., 1998; Nisbet
81 et al., 2014). ~~However~~ Subsequently, a low growth rate was sustained over 1999 to
82 2006, except for ~~special~~ years (2002/2003) with El Niño events (2002/2003)
83 (Dlugokencky et al., 1998). The annual growth rates dropped from ~12 ppb yr⁻¹ to near

84 0 from ~~the~~ late 1980s to ~~1999–2006~~ (Nisbet et al., 2019), followed by a renewed CH₄
85 ~~increase since then~~. ~~But thereafter, the atmospheric CH₄ concentration keeps rising~~
86 ~~from 2007~~ (Nisbet et al., 2019). During 2007–2013, the annual growth rate of methane
87 was 5.7 ± 1.2 ppb yr⁻¹. After 2013, the atmospheric CH₄ grew ~~even~~ at rates not observed
88 ~~after 1980s~~, such as 12.7 ± 0.5 ppb yr⁻¹ in 2014 and 10.1 ± 0.7 ppb yr⁻¹ in 2015 (Nisbet
89 et al., 2016, 2019). ~~In the most recent decade, t~~The overall global mean growth rate was
90 7.1 ppb yr⁻¹ ~~in recent 10 years~~ (WMO, 2019). ~~And the not expected increase since 2007~~
91 ~~would make it difficult to meet the targets of carbon emission reduction in the future.~~
92 The World Meteorological Organization/Global Atmospheric Watch ~~P~~programme
93 (WMO/GAW) annual greenhouse gas bulletin revealed that ~~the~~ globally averaged CH₄
94 mole fraction of CH₄ reached a new high with 1869 ± 2 ppb in 2018 (Rubino et al.,
95 2019), which was ~259% of pre-industrial levels (~722 ppb around 1750 C.E.
96 (Common Era)) (Etheridge et al., 1998; WMO, 2019).

97 ~~The a~~Atmospheric CH₄ is mainly emitted from natural sources (e.g. about 40%
98 ~~from, e.g.,~~ ruminants and wetlands) and anthropogenic sources (e.g. about 60% ~~from,~~
99 ~~e.g.,~~ paddies, cattle ranches, coal mines, fossil fuels, and biomass burning) (Hausmann
100 et al., 2016; Saunio et al., 2016). ~~ObservationsStudies~~ from ~~the~~ GAW ~~observations~~
101 indicated that the causes of ~~the~~ recent increase were likely attributed to anthropogenic
102 emissions at mid-latitudes in the ~~N~~orthern ~~H~~emisphere and the wetlands in the tropics
103 (WMO, 2019). The rapid development of population growth, economic expansion, and
104 ~~countries~~ urbanization of countries has led to more and more fossil fuel production and
105 consumption (e.g.e.g., the large-scale exploitation of natural gas, oil, and coal) and
106 biomass burning, ~~C~~onsequently, large amounts of anthropogenic CH₄ were emitted
107 around the world in recent years (Galloway, 1989; Streets and Waldhoff, 2000; Wang
108 et al., 2002; Lin et al., 2014; Hausmann et al., 2016). The recent carbon isotope study
109 revealed that biogenic emissions might also have driven CH₄ increases, including
110 microbial sources ~~whether~~ from rice, paddies, ruminants, termites, enteric fermentation,
111 or a combination of these (Nisbet et al., 2016; Schaefer et al., 2016; Wolf et al., 2017).

112 ~~About~~ 90% of ~~the~~ CH₄ destruction in the atmosphere ~~is~~ ~~are~~ ~~mainly~~ from the reaction
113 with hydroxyl radicals ($\cdot\text{OH}$) (Vaghjiani and Ravishankara, 1991; Bousquet et al.,
114 2011), an important oxidant in the troposphere (Logan et al., 1981). Therefore, the inter-
115 annual ~~variability~~ ~~variation~~ of $\cdot\text{OH}$ or the decline ~~of the~~ oxidative capacity of the
116 atmosphere may also cause the recently increase ~~in the~~ CH₄ growth rates (Rigby et al.,
117 2017; Turner et al., 2017). ~~Even though, the exact causes of significantly increase in~~
118 ~~the~~ CH₄ emissions in past years ~~are still remained unclear and debated, especially for~~
119 ~~the anomalous periods with a suddenly large growth, due to the sparse time and space~~
120 ~~sparsity of measurements and the crude model approaches, which limited our~~
121 ~~understanding of the global variation of atmospheric CH₄ (Saunois et al., 2019; Weber~~
122 ~~et al., 2019). As long as the reasons for rising CH₄ emissions contributed by natural~~
123 ~~sources (e.g. wetlands), anthropogenic sources (e.g. fossil fuels), or climate change~~
124 ~~feedbacks remain uncertainties~~ a consequence, it will be impractical to predict CH₄
125 trends in the future, and then to develop realistic management (Nisbet et al., 2019).
126 Therefore, it is essential to establish typical observing regions and perform long
127 ~~observations.~~

128 ~~Systematic observations are a prerequisite to~~ To get an accurate understanding of
129 ~~the spatial and temporal behaviors of~~ atmospheric CH₄, ~~concentrations.~~ a systematic
130 ~~observations network would perform the best. Hundreds of CH₄ observation stations~~
131 ~~worldwide are running under the framework of WMO/GAW.~~ Since 1978, systematic
132 measurements of atmospheric CH₄ ~~have been taken~~ began around the world (Blake et
133 al., 1982; Rasmussen and Khalil, 1984; Dlugokencky et al., 1994). On the northern
134 slope of the Mauna Loa volcano, Hawaii, there exists the first global station, Mauna
135 Loa (MLO), which ~~h~~was performed ~~at~~ about 3397m above sea level (a.s.l.) and far away
136 from local sources and sinks. It has the longest records of continuous atmospheric CH₄
137 observation (Keeling et al., 1976). ~~Recently~~ Later, ~~many other types of the sites~~ stations
138 ~~have been~~ were installed for CH₄ observation system, ~~such as~~ such as the Barrow
139 (BRW), South Pole (SPO) (polar site, ~~SPO~~) (Dlugokencky et al., 1995), Cape Grim

140 (CGO) and Minamitorishima (MNM) in Australia (coastal/island sites) (Pearman and
141 Beardsmore, 1984; Wada et al., 2007), Minamitorishima (MNM) in Japan
142 (coastal/island sites) (Wada et al., 2007), Jungfraujoch (JFJ) in Switzerland and Mount
143 Waliguan in China (continental mountain site) (Zhou et al., 2004; Loov et al., 2008).
144 Hundreds of CH₄ observation stations worldwide are currently running under the
145 framework of the WMO/GAW. ~~Even though, the exact causes of significantly increased~~
146 ~~CH₄ emissions in past years are still remained unclear and debated, especially for the~~
147 ~~anomalous periods with suddenly large growth, due to the time and space sparsity of~~
148 ~~measurements and the crude model approaches, which limited our understanding of the~~
149 ~~global variation of atmospheric CH₄ (Saunois et al., 2019; Weber et al., 2019). As long~~
150 ~~as the reasons for rising CH₄ emissions contributed by natural sources (e.g. wetlands),~~
151 ~~anthropogenic sources (e.g. fossil fuels), or climate change feedbacks remain~~
152 ~~uncertainties, it will be impractical to predict CH₄ trends in the future, and then to~~
153 ~~develop realistic management (Nisbet et al., 2019). Therefore, it is essential to establish~~
154 ~~typical observing regions and perform long observations.~~

155 China has the largest anthropogenic CH₄ emissions in the world (Janssens-
156 Maenhout et al., 2019). ~~The~~ Qinghai-Tibetan Plateau with has an average altitude of
157 over 4000m a.s.l., which has long been recognized as the roof of the world. By
158 coincidence, the two largest CH₄ source regions in the world (i.e.i.e. Eastern China and
159 Northern India) are trapped by the Tibetan Plateau in betweenthe middle (Zhang et al.,
160 2011; Fu et al., 2012; Wilson and Smith, 2015). Under the characteristics of special
161 geographical conditions, lower population density, rarely industrial activities, and high
162 sensitivity to external disturbances, the Tibetan Plateau is undoubtedly one of the ideal
163 regions to observe a continual CH₄ signal (Zhou et al., 2005; Fu et al., 2012; Zhang et
164 al., 2013). Most of the previous studies reported on the short-term CH₄ variations in
165 China and concluded that the importance of long-term observation ~~(Cai et al., 2000;~~
166 ~~Zou et al., 2005; Wang et al., 2009; Fang et al., 2013),~~ which is of great value to enhance
167 the understanding of the global carbon cycle ~~(Cai et al., 2000; Zou et al., 2005; Wang~~

168 [et al., 2009; Fang et al., 2013](#); Yuan et al., 2019). As the rapid development of China
169 and India continues, the ~~year to year difference as well as the sources and~~
170 ~~sinks~~characteristics of CH₄ on the Tibetan Plateau, might change significantly over
171 time. Since 1994, in-situ measurements of atmospheric CH₄ have been launched at the
172 Mt. Waliguan (WLG) station. To study the long-term variations of the atmospheric CH₄
173 ~~at the WLG~~ and get a new insight ~~of on~~ its characteristics in the inland of ~~the~~ Eurasia,
174 ~~in this study, we evaluated~~ the performance of a ~~24-26-year-year~~long term of in situ
175 observations ~~of CH₄~~ at the Mt. Waliguan baseline observatory were evaluated, which is
176 the longest ~~time observing~~ records in China. Temporal patterns, annual variations, long-
177 term trends, air mass transports, and the spatial distribution of potential sources were
178 analyzed. In addition, the case studies combining atmospheric carbon monoxide (CO)
179 measurements and a separate analysis between the Tibetan Plateau and the city regions
180 were performed to constrain the contribution of anthropogenic emissions.

181

182 2 Methodology

183 2.1 Measurement site

184 The Mt. Waliguan (WLG, 36.28° N, 100.09° E, 3816m a.s.l.) station is situated at the
185 edge of northeastern ~~of the~~ Tibetan (Qinghai-Xizang) Plateau, which ~~was is~~ in remote
186 western China and isolated from populated and industrial regions ([Fig. 1](#)~~Fig. 1~~). WLG
187 ~~was is~~ the only WMO/GAW global background station in Eurasia and is running-run
188 by the China Meteorological Administration (CMA). The surrounding ~~areas~~ of the site
189 are pristine with sparse vegetation, naturally arid and semi-arid grasslands. Small farms
190 with yak and sheep are in the valley. Two adjacent large cities Xining (~2.2 million
191 populations) and Lanzhou, are located about 90 km northeast and 260 km east of the
192 station, respectively. The Longyangxia hydroelectric station (~380 km²) is located
193 approximately 13 km south to southwest of ~~the the~~ WLG. The predominant winds at
194 the WLG are mainly from southwest and east in winter and summer, respectively (Zhou

195 et al., 2004; Zhang et al., 2011), which is controlled by Tibetan Plateau monsoon.
196 Simultaneously, diurnal variations of vertical winds at the WLG ~~is~~ are influenced by
197 mountain-valley breezes, where upslope flow brings heated air masses from the
198 boundary layer to the site in daytime and downslope flow results in cool air masses
199 transport from the mountain peak to the site. Under this unique location, the observation
200 at the WLG ~~could~~ can obtain essential information on CH₄ sources and sinks from
201 Eurasia (Zhou et al., 2005; Zhang et al., 2013).

202

203 2.2 Instrumental setup

204 Atmospheric CH₄ has been measured quasi-continuously using a HP 5890 gas
205 chromatograph (GC) equipped with a flame ionization detector (FID) since July 1994,
206 and an Agilent 6890N GC equipped with a FID since June 2008. Both of the systems
207 used the same sampling procedures. A Cavity Ring Down Spectroscopy system (Picarro
208 G1301) began in January 2009 and the instrument was upgraded to Picarro G2401 in
209 2015. Ambient air is delivered to the above systems at about 5 L/min by a KNF
210 Neuberger N2202 vacuum pump via a dedicated 0.95 cm o.d. sample line from an 80m
211 intake line attached to an 89m steel triangular tower located approximately 15m from
212 the main observatory. The residence time of the ambient air from the top of the tower
213 to the instrument is 30 s. The ambient air is first passed through a 7 mm stainless steel
214 membrane filter located upstream of the pump and then (after the pump) passed through
215 a pressure relief valve set at 1 atm to release excess air and pressure. The ambient air is
216 then dried to a dew point of approximately -60°C by passing it through a glass trap
217 submerged in a -70°C methanol bath. All standard gases supplied to the instruments are
218 from pressurized 37.5 L treated aluminum alloy cylinders fitted with high-purity, two-
219 stage gas regulators. Stainless steel tubing (0.32 cm o.d., 0.22 cm i.d.) is used for the
220 standard gas sample line and the ambient sample line after the cold trap. ~~The An~~
221 automated sampling module equipped with a VICI 8 ports valve is designed to sample
222 from separate gas streams (standard tanks and ambient air). According to the

223 comparability target of WMO/GAW program (WMO, 2019), methane mole fractions
224 are referenced to a Working High standard (WH) and a Working Low standard (WL).
225 Additionally, a calibrated cylinder filled with compressed ambient air is used as a Target
226 gas (T) to check the precision and stability of the system routinely. Diagram of the
227 observing system during different periods could be seen at Zhou et al. (2004) and Fang
228 et al. (2013). Here, we focus on the longest continuous measurements of CH₄ from
229 August 1994 to ~~May-December~~ 2019~~7~~ at WLG. Data gaps in limited periods are
230 because of the malfunction of instrument and the maintenance of the sampling system.

231 The records of CO in this study was initially observed by an RGA-3 gas
232 chromatograph (GC) equipped with an HgO reduction detector (Trace Analytical Inc.)
233 since 1994. An automated sampling module was designed to sample from ambient air
234 and a series 9 standards. Detailed diagram of the system was described by Zhang et al.
235 (2011). Since 2010, the CO has been measured by the Cavity Ring Down Spectroscopy
236 instrument (Picarro G1302 and G2401 since 2015). The scale for all of the CO
237 measurement were further updated to WMO X2014A.

238

239 2.3 Data processing

240 Most on-site CH₄ observations were unavoidably influenced by local sources and other
241 complex conditions (~~e.g., e.g.~~ traffic transportation, various topography). As a result, the
242 records cannot fully represent the regional atmospheric CH₄ in well-mixed conditions
243 (Liu et al., 2019). To ~~precisely~~ get regionally representative ~~regional~~ records, we
244 excluded the CH₄ ~~measurements data~~ influenced by local sources adjacent to the site
245 (~~e.g., e.g.~~ agricultural fields, cities, traffic emissions). The hourly CH₄ data were
246 classified as Locally/Regionally representative events through the meteorological
247 approach, which was based on essential meteorological information, similar to previous
248 studies by Zhou et al. (2004) and Liu et al. (2019). In this study, the CH₄ records
249 associated with local surface winds from selected sectors, (~~i.e., i.e.~~ NNE-...-ENE in
250 spring, NE-...-SE in summer, NE-...-ESE in autumn, and NE-ENE in winter,) were

251 flagged as locally ~~representative~~influenced (27.0%). Subsequently, we ~~further~~ rejected
252 portion of daytime records to minimize the effect of human activities (16.9%), ~~e.g., e.g.~~
253 ~~rush hours~~), including 9:00-13:00 LT (local time) in ~~spring and~~ summer, ~~9:00-14:00 LT~~
254 ~~in autumn, and~~ 10:00-17:00 LT in winter. Finally, we filtered CH₄ data into locally
255 ~~events-influenced~~ when the surface wind speeds ~~were~~as less than 1.5 m s⁻¹ to minimize
256 the very local accumulation (9.2%).

257 In order to further investigate the influence of anthropogenic emissions from cities
258 and remote area as Tibetan plateau, investigate the characteristics of atmospheric CH₄,
259 we divided the CH₄ observations data into two main influencing regions according to
260 the above analysis, including the geographical conditions, the effect of surface winds,
261 the long-range transports, and the potential source distributions. The first region was
262 covered the northeast and to southeast (NNE-...SE) of the WLG, which was denoted
263 as City Regions (CR). The second region was located the south to west (S-...-W) of
264 the station and was well known Tibetan (Qinghai-Xizang) Plateau (TP) (Fig. S21).
265 Accordingly, the hourly CH₄ records when the surface winds coming from these sectors
266 were divided into two subsets (i.e. i.e. TP and CR). The long-term variations between
267 the two regions as well as the total regional time series (Total) were further compared
268 and analyzed to explore new sight of atmospheric CH₄ variation at WLG.

270 To understand the influence of local surface wind, ~~the~~ hourly CH₄ data ~~was~~ was
271 ~~calculated binned versus into~~ 16 horizontal wind direction classes (Fang et al., 2013).
272 In this study, we used the ‘polarPlot’ function ~~located~~ in the ‘openair’ package of the
273 statistical software R (R Core Team, 2020). It shows the bivariate (~~i.e. i.e.~~ wind speed
274 and wind direction) polar plot of CH₄ concentrations, and the concentrations are
275 calculated as a continuous surface by modelling using smoothing techniques (Carslaw
276 et al., 2006; Diederich, 2007). Also, conditional probability function (CPF) was used to
277 ~~deteet~~ investigate the probability of which wind directions are dominated by high CH₄
278 mole fractions (Uria-Tellaetxe and Carslaw, 2014). In order to study the pollution

279 transport pathways of air masses at WLG, the cluster analysis of 3 days back trajectories
 280 was applied using the Hybrid Single-Particle Lagrangian Integrated Trajectory
 281 (HYSPLIT) dispersion model (Draxier and Hess, 1998; Rousseau et al., 2004) on the
 282 strength of gridded meteorological data (2004-2017) from the National Oceanographic
 283 and Atmospheric Administration’s Air Resources Laboratory (NOAA-ARL). The
 284 spatial resolution of the model is 0.5×0.5 degree and the model height is 10 km a.s.l.
 285 We computed the back trajectories coincident with hourly CH₄ mole fractions. The
 286 trajectories ~~_in four months, including for~~ January, April, July and October, were
 287 calculated to represent the seasons of winter, spring, summer and autumn, respectively.
 288 The spatial source distributions of annual CH₄ were analyzed using the Potential Source
 289 Contribution Function (PSCF) approach, which computed the conditional probability
 290 of the residence times of air parcels with greater concentration than threshold transport
 291 to the exact~~ly~~ receptor site (Ashbaugh et al., 1985). In this study, PSCF value was
 292 calculated in 0.5×0.5-degree grid cell (*i, j*):

$$293 \quad \text{PSCF}_{ij} = m_{ij} / n_{ij} \quad (1)$$

294 n_{ij} represents the number of endpoints that terminate in the *ij*th grid cell, while the
 295 number of trajectories with concentration exceed the threshold value was defined as m_{ij}
 296 (Polissar et al., 1999). In order to reduce the abnormal influence of small n_{ij} values in
 297 some grid cells, PSCF_{*ij*} was further computed by an arbitrary weighting function W_{ij} as
 298 below.

$$299 \quad W_{ij} = \left\{ \begin{array}{ll} 1.00 & 3n_{ave} < n_{ij} \\ 0.70 & 1.5n_{ave} < n_{ij} \leq 3n_{ave} \\ 0.42 & n_{ave} < n_{ij} \leq 1.5n_{ave} \\ 0.05 & n_{ij} \leq n_{ave} \end{array} \right\} \quad (2)$$

300 W_{ij} represents the weight of cell (*i, j*), n_{ij} is the number of trajectory endpoints that fall
 301 in the *ij*th grid cell, while the n_{ave} shows the mean number of the endpoints in all grid
 302 cells.

303 In order to fill the data gaps so as to evaluate the long-term CH₄ trend, we applied
 304 the curve fitting approach by Thoning et al. (1989). We also calculated the trend curve

305 that excluded the influence of seasonal variation, and then got the annual growth rates
306 ~~of by~~ the average of the first derivative of the trend curve. The function consists of the
307 polynomial part and the annual harmonics part:

$$308 \quad f(t) = a_0 + a_1 t + a_2 t^2 + \dots + a_{(k-1)} t^{(k-1)} + \sum_{n=1}^{nh} c_n [\sin(2n \pi t) + \varphi_n] \quad (3)$$

309 'k' represents the number of polynomial part. 'nh' is the number of harmonics part. We
310 applied k = 3 polynomial terms (a quadratic) for multi-year trends and nh = 4 yearly
311 harmonics for seasonal cycles in this study. The fast Fourier transform (FFT) was
312 utilized to smooth the fitting residuals (Press et al., 1992).

313 The significant difference test was applied by the 'scheirerRayHare' function in
314 the 'rcompanion' package of R software, which is a non-parametric test for two-way
315 ANOVA analysis. ~~And t~~The multiple comparison was used by Wilcoxon rank sum test
316 with R (R Core Team, 2020). For the correlation analysis between CH₄ and CO, we
317 obtained the detrended time series of CH₄ and CO from 2004-2017 based on the method
318 by Thoning et al. (1989). The detrended values are denoted as ΔCH_4 and ΔCO , which
319 are the original data points minus the trend curve. ~~The detrended values are denoted as~~
320 ~~ΔCH_4 and ΔCO .~~ To accurately obtain the correlation slopes of ΔCO ~~$-\Delta\text{CH}_4$~~ and ΔCH_4
321 ΔCO , i.e. $\Delta\text{CO}/\Delta\text{CH}_4$ ~~accurately~~, a rolling linear regression was applied to the ΔCH_4
322 and ΔCO time series by the 'roll_lm' function in 'roll' package of R (R Core Team,
323 2020). We successively moved a 24-h time window by 1 h over the whole time series.
324 Similar to the study by Tohjima et al. (2014), we set ~~three~~3 criteria to achieve a better
325 quality control of the slopes. When (i) the number of CH₄ record is less than ~~five~~5 in
326 24 hours, (ii) the coefficient variation of the correlation slope is more than 15%, (iii)
327 the absolute value of the correlation slope is less than 0.8 ($|R| < 0.8$), the correlation
328 slopes were identified as statistically insignificant and inaccurately and were rejected.
329 In order to understand the year to year variations, we ~~further~~ analyzed the CH₄ variation
330 over different periods ~~over in~~ 1994-20197. The entire CH₄ time series were divided into
331 five observing periods, i.e. 1994-1997, 1998-2002, 2003-2007, 2008-2012, and 2013-
332 2019, according to the significant stages or the critical time period of atmospheric CH₄

333 variations from ~~previously~~ studies (Zhou et al., 2004; Fang et al., 2013; Zhang et al.,
334 2013; Nisbet et al., 2019; WMO, 2020), ~~appropriately every five years a period~~. Unless
335 special notes, the average values in this study ~~were~~ are presented with 95% confidence
336 intervals (CIs).

338 3 Results and discussion

339 3.1 Extracting the regional atmospheric methane

340 To precisely understand the characteristics of atmospheric CH₄, including seasonal
341 cycles and long-term trends, it is vital to identify the CH₄ records in well-mixed air
342 without local contaminations (Liu et al., 2019). In this study, hourly CH₄ measurements
343 between 1994 and 2019 were analyzed, resulting in 64.0% of the CH₄ data being
344 classified as regionally representative, with an average CH₄ mole fraction of 1865.8 ±
345 0.4 ppb. The average of the locally influenced data (1868.2 ± 0.3 ppb) was larger than
346 that of the regionally representative records (Table 1). The filtered regional/local time
347 series was shown in Figure 2. It can be seen that the CH₄ mole fractions increased from
348 1994 to 2019. The atmospheric CH₄ showed a strong growth and displayed large
349 fluctuation. In 1994, the average CH₄ mole fraction was only 1799.0 ± 0.4 ppb,
350 however, the average value increased 133 ppb by the year 2019 (1932.0 ± 0.1 ppb)
351 (Table 2). Compared with the global average mole fractions in recent years, i.e. 1853 ±
352 2 ppb in 2016, 1859 ± 2 ppb in 2017, and 1869 ± 2 ppb in 2018 (WMO, 2019; 2020),
353 the atmospheric CH₄ mole fractions at the WLG were significantly higher. These results
354 indicated that the WLG was affected by strong CH₄ sources in recent years, which were
355 possibly due to the influence of the two largest source regions of Northern India and
356 Eastern China (Fang et al., 2013; Zhou et al., 2004).

357 2.43.2 Diurnal variations

359 ~~Generally,~~ Distinct diurnal cycles were observed in four seasons during 1994-

360 ~~2017~~2019 at the WLG. The CH₄ mole fraction increased from early morning, ~~and~~
361 reached ~~at~~ the maximum at noon, and ~~had~~ a trough in ~~the~~ late afternoon (Fig. 3Fig. 2f).
362 However, differences also existed ~~in~~from ~~different~~ ~~four~~ seasons. In spring ~~and~~ summer,
363 the atmospheric CH₄ ~~apparently~~ increased from 9:00 to 13:00 LT ~~at~~ noon, with ~~at~~ the
364 daily amplitude of ~~5.87 ± 2.84~~ and ~~4.4 ± 3.4~~ ppb, respectively ~~ppb~~(Table S1). ~~In~~ summer,
365 ~~the elevated CH₄ also appeared during 9:00–13:00 LT at noon, and the daily amplitude~~
366 ~~was 4.3 ± 2.6 ppb.~~ In autumn, the diurnal variation showed ~~an~~the ~~mean~~-amplitude of
367 ~~4.35 ± 32.14~~ ppb, ~~significantly elevated CH₄ reached at 9:00–13:00 LT~~ with one peak
368 at noon. In winter, ~~a~~ ~~largely~~ increase of CH₄ ~~sing~~ ~~pr~~ was found ~~during~~ ~~esented~~ in the
369 ~~daytime at 9:00–17:00 LT, with the largest largest peak to trough amplitude of 7.1 ± 2.9~~
370 ~~ppb of 6.2 ± 2.4 ppb among four seasons.~~ For the diurnal variation over the whole
371 ~~monitoring period, the highest CH₄ mole fraction was observed in winter and the~~
372 ~~minimum value was found in spring (Fig. 2f).~~

373 Different patterns for diurnal CH₄ cycles were also found over different periods. In
374 1994–1997 ~~and~~ 1998–2002, the CH₄ mole fractions in ~~the~~ winter were ~~apparently~~ higher
375 than the other seasons (Fig. 3Fig. 2a-b), ~~which~~ ~~was~~ likely due to the manner of heating
376 ~~(e.g. large biomass burning) as well as the adverse diffusion conditions in cold weather~~
377 ~~(Fang et al., 2013).~~ However, ~~But its value was~~ the highest ~~value was found~~ in summer
378 during the periods of 2003–2007, 2008–2012, and 2013–~~2017~~2019 (Fig. 3Fig. 2c-e),
379 ~~which could be ascribed to the transport of anthropogenic emissions (e.g. southeast~~
380 ~~cities) by the meteorological factors (Zhang et al., 2013).~~ Additionally, in summer, the
381 ~~intense herding or grazing activities around the WLG have enhanced the regional CH₄~~
382 ~~emissions and hence contributed to the higher CH₄ mole fractions (Zhou et al.,~~
383 ~~2004).~~ The atmospheric CH₄ values in winter were gradually falling behind the other
384 ~~seasons, and the gaps among different seasons were increasing, especially for summer.~~

385 ~~Similar to the previous studies by Zhang et al. (2013) and Fang et al. (2013), Before~~
386 ~~2002, diurnal cycles in four seasons were ambiguous before 2002 (Fig. 3Fig. 2a-b), but~~
387 ~~significant diurnal variations appeared afterwards (Fig. 3Fig. 2c-e), which indicated~~

388 that the local sources were weak at the WLG in the past. The apparent diurnal cycles
389 after 2002 may be attributed to the intense activities by humans (e.g. grazing, burning
390 fuel), which were enhanced in the daytime and weakened at nighttime (Fang et al.,
391 2013). The peak to trough amplitude almost increased ~~along with the time~~ in almost all
392 seasons. For example, in spring, the amplitude was 6.5 ± 3.01 , 4.7 ± 2.14 , 5.6 ± 2.76 ,
393 76.02 ± 32.14 and 6.98 ± 3.14 ppb over the periods of 1994-1997, 1998-2002, 2003-
394 2007, 2008-2012, and 2013-~~2017~~2019, respectively (Table S1~~Table S1~~). The
395 meteorological conditions (e.g. diffusion and transport) could have also contributed to
396 the increasing CH₄ amplitude. The WLG was remote from the populated center,
397 therefore, the good diffusion conditions in the daytime may have brought high
398 anthropogenic emitted CH₄ to the site. The increasing amplitude and CH₄ mole
399 fractions suggested that the WLG was increasingly affected by local and regional
400 anthropogenic sources, such as gas exploitation, and grazing (Zhou et al., 2004).

403 2.53.3 The impact of local surface winds

404 ~~As observed by~~Similar to the ~~previous~~ short-term study variations by Zhou et al. (2004),
405 the atmospheric CH₄ at the WLG was significantly influenced by local surface winds
406 from the northeast ~~and to~~ southeast sectors (Fig. 4~~Fig. 3f~~). Slight differences were also
407 found among seasons. In the spring, when the wind was originating from the NNE-NE-
408 ENE-E sectors, the atmospheric CH₄ was enhanced by 2.5-6.5 ppb as compared to the
409 seasonal average (1839.7 ± 1.4 ppb) ~~when the wind was originating from NNE-NE-~~
410 ~~ENE-E sectors.~~ In the summer and autumn, the wind from the NNE-NE-ENE-E-ESE
411 produced higher CH₄ mole fractions, with an enhancement of ~~about~~ 3-9.5 ppb and 4
412 ~~to~~ 18 ppb, respectively. In the winter, similar to that in spring, the CH₄ mole fractions
413 significantly ~~increased~~ elevated ~~from the~~ from NNE-...-E sectors ~~wind sectors that same~~
414 ~~as those found in spring,~~ with a value of 7-21 ppb higher than the seasonal average
415 (1854.5 ± 4.8 ppb). In summer, the prevailing winds were from the NE-...-ESE sectors

416 (~46%) (Fig. S2), and the CH₄ mole fractions were also higher in the related sectors.
417 However, in autumn and winter, although the prevailing wind and the high wind speed
418 were from SSW-...-W sectors (~ 40-50%) (Fig. S2), the high CH₄ mole fractions were
419 from the opposite wind sectors of NNE-....-ESE (Fig. 4f), suggesting that strong local
420 sources were distributed from northeast to southeast of the WLG (city regions), and
421 even covered the emission of natural sources. Causes of the elevated CH₄ from these
422 sectors could be attributed to the large plantations of highland barley as well as the high
423 population density in those areas (Fang et al., 2013). The two largest cities of Xining
424 (with population of ~2.2 million) and Lanzhou (with population of ~4 million) are also
425 situated in the northeast and east of the WLG, respectively. The heavy human activities
426 from anthropogenic fossil combustion, landfills, and livestock, could have also emitted
427 large amounts of CH₄. Based on the data from the Emissions Database for Global
428 Atmospheric Research (EDGAR), the increase of CH₄ emission was 500 kg yr⁻¹ in these
429 two regions throughout 1994-2015 (Crippa et al., 2019). Also, the Yellow River Canyon
430 industrial area (YRC), which is ~500 km northeast of the WLG, may also have
431 contributed to the high CH₄ values (Zhou et al., 2003). With the rapid development of
432 land-use, water utilization, and agriculture sources in the YRC, large CH₄ emissions
433 could have easily transported to the WLG. Previous studies on black carbon (BC) and
434 carbon monoxide (CO) also revealed that the high CH₄ values at the WLG in winter
435 were a result of transport from the YRC (Tang et al., 1999; Zhou et al., 2003).
436 The wind-rose distribution of CH₄ mole fractions indicated that the ~~Relatively, the~~
437 ~~amplitude of enhancements in winter and autumn were larger than those in spring and~~
438 ~~summer.~~

439 ~~What interesting is that wind sectors~~ elevated CH₄ mole fractions varied in
440 ~~different periods. The~~ early periods (i.e.i.e. 1994-1997 and 1998-2002) ~~to~~ were different
441 ~~from the~~ recent periods (2003-2007, 2008-2012, and 2013-2017). The elevated CH₄
442 was predominately from ~~the~~ about ENE-E-ESE-SE-~~SE~~-SSE sectors in ~~the~~ early years (Fig.
443 ~~4~~Fig. 3a-b), but evolved to ~~the~~ NNE-NE-ENE-E sectors in later years (Fig. ~~4~~Fig. 3c-e).

444 Furthermore, the amplitude of enhancements was ~~almost also~~ increasing ~~continuously~~
445 along with the progression of the time. For example, in autumn, the maximum ~~CH₄~~
446 ~~mole fractions~~ were from the E in 1994-1997, ENE in 1998-2002, ENE again in 2003-
447 2007, NE in 2008-2012, and finally ENE in ~~sectors in 1994-1997, 1998-2002, 2003-~~
448 2007, 2008-2012 and 2013-2017, with the successively ~~increasing~~ enhancements of
449 8.6, 12.1, 14.7, 16.8, and 19.7 ppb, respectively. Therefore, the local surface wind from
450 the city regions had an increasing effect on the atmospheric CH₄ at the WLG.

451 ~~We~~ The CPF was applied ~~the CPF~~ to hourly CH₄ and CO data by considering
452 intervals of entire data percentiles including 0-20th, 20th-40th, 40th-60th, 60th-80th, and
453 80th-100th ~~0-20, 20-40, 40-60, 60-80 and 80-100~~ to draw the CPF polar plot (Uria-
454 Tellaetxe and Carslaw, 2014). It ~~was~~ clear that the different sources only affected the
455 CH₄ mole fractions ~~in~~ different percentile ranges (Fig. 5), meaning that the ~~specific~~
456 sources were prominent for specific percentile ranges. ~~For example,~~ for most wind
457 speed-directions, the CPF probability of the CH₄ being greater than the 60th percentile
458 was ~~trending to zero~~ (Fig. 4). ~~And it is apparent that,~~ and most sources contributed to
459 ~~the less than 60th percentiles~~ of less than the 60th for the CH₄ mole fraction (e.g.e.g.
460 40th-60th 40-60) (Fig. 4). These results implied that most areas around the WLG had a
461 small contribution to the CH₄ emissions. In addition, ~~The specific sources were~~
462 prominent for specific percentile ranges. ~~The~~ wind from the southwest and the
463 southeast was important ~~for~~ the ~~the~~ higher percentiles, resulting in the
464 highest CH₄ mole fractions of 1849-1872 ppb for the 60th-80th 60-80 percentiles and
465 1872-2031 ppb for ~~80-100~~ the 80th-100th percentiles (Fig. 5Fig-4), which revealed the
466 existence of a southeastern and a southwestern strong source region. The anthropogenic
467 emissions from cities (e.g. Lanzhou or Chengdu) were the only cause for high values
468 in the southeast, and the southwest region that was farther away from the WLG was
469 possibly due to sources from other countries, such as India. ~~It's's more obvious~~
470 that the CO sources ~~the CO showed~~ gradually shifted ~~sources~~ with the increase of the
471 percentile ranges (Fig. 5). The areas where the CPF probabilities were higher ~~is~~ were

472 to the NW-SW sectors when ~~the~~ percentages ranged from 0 to 40th. Nevertheless, when
473 ~~the percentages were larger than data exceeded the 60th 60percentileth~~, the high
474 probability areas completely moved to NE-SE sectors (~~Fig. 4~~). Unlike that of CH₄, the
475 high CO mole fractions were consistent from the east regions (urbanized areas) (Fig.
476 5), indicating strong anthropogenic sources in the city regions (e.g. Xining and
477 Lanzhou) and different source distributions between CH₄ and CO around the WLG
478 (Zhang et al., 2011).

479

480 2.6.3.4 Air mass pathways and Long-range transport and potential source 481 distributions

482 2.6.13.4.1 Air mass transports

483 Figure ~~5-6~~ illustrates the cluster analysis ~~to~~ of the 3-day back trajectories ~~to WLG~~
484 ~~betweenduring~~ 2004 ~~and~~ -2017. In ~~the~~ spring, ~~the~~ majority of the air masses were from
485 ~~the~~ west and northwest regions, which accounted for about 24% (cluster 3) and 44%
486 (cluster 5) of ~~the~~ total trajectories (~~Fig. 6~~Fig. 5a). These air masses were also
487 accompanied ~~with~~ ~~by~~ higher CH₄ mole fractions than those from ~~the~~ east and northeast
488 regions, ~~such as~~ ~~i.e.~~ cluster 1 (13.3% of total) and cluster 4 (11.69%) (~~Table 3~~Table 4).
489 The largest enhancement was ~~in cluster 3 at~~ ~18 ppb (relative to ~~the spring~~ average) ~~by~~
490 ~~cluster 3~~. In ~~the~~ summer, 45% of ~~the~~ air masses (~~cluster 1~~) were from eastern regions
491 (~~cluster 1~~). ~~However,~~ ~~But~~ the highest CH₄ mole fractions were ~~in~~ on cluster 2 and cluster
492 5 from northwest and west regions, ~~respectively,~~ ~~However,~~ ~~although~~ low percentages
493 ~~were found in each of those clusters, i.e. 26% in were found~~ (cluster 2 and 7%: 26%: ~~in~~
494 cluster 5: 7%) (~~Fig. 6~~Fig. 5b). ~~The highest CH₄ mole fraction was associated with~~
495 ~~cluster 2, with ~9 ppb larger than the average in summer.~~ In ~~the~~ autumn, ~~a~~ large
496 proportion of air masses ~~cluster 2 (49%) and cluster 3 (32%) was~~ ~~originated~~ ~~from~~
497 ~~the~~ west and southwest stations, ~~such as~~ ~~respectively~~ cluster 2 (49%) and cluster 3 (32%)
498 (~~Fig. 6~~Fig. 5c). ~~C~~The highest CH₄ ~~was from~~ cluster 3 ~~showed the highest CH₄ mole~~

499 ~~fraction, with an~~with enhancement of ~4 ppb ~~relative to~~than the seasonal average.
500 ~~Similar to autumn~~In winter, the air masses ~~were~~primarily ~~came~~from northwest ~~and~~
501 ~~southwest regions in winter, including e.g. northwest~~cluster 3 (59%), ~~and southwest~~
502 cluster 1 (34%) ~~regions (Fig. 6Fig. 5d)- and t~~The ~~cluster 1 brought the~~highest CH₄ mole
503 fractions ~~was on cluster 1~~with ~~an~~the enhancement of ~7 ppb ~~over~~than the ~~average~~
504 ~~values~~seasonal average.

505
506 The air masses from east and northeast regions passed over the cities of Xining and
507 Lanzhou (capital of Gansu province), which are the populated centers and industrial
508 areas (Fig. 6). However, the higher CH₄ mole fractions were not observed when the air
509 masses were from these sectors, the high values were frequently brought by air masses
510 that came from the northwest to southwest (Table 3). The higher values seen in the
511 northwest to southwest airflow were because the air masses had passed through the
512 northwest of the Qinghai province and the central area of the Xinjiang Uygur
513 Autonomous Region (XUAR). This is where the Ge'ermu urban area (the second
514 largest city of Qinghai) was located and where there was with rapid industrial
515 development, natural gas and petroleum resource exploitation, and residue burning of
516 large crops, hence the CH₄ emissions were strong (Fang et al., 2013; Zhang et al., 2013).
517 This result was also similar to the CPF percentile analysis (Fig. 5), as time went by, the
518 southwest or northwest regions that were farther away from the site became the
519 strongest source regions to the WLG.

520 521 2.6.23.4.2 **Spatial distribution of potential source regions**

522 ~~In this study, t~~The potential sources were analyzed over different periods, ~~i.e.i.e.~~ 2004-
523 2007, 2008-2012, ~~and~~ 2013-2017 (Fig. 7). Generally, the ~~strongest potential~~sources
524 were located ~~at the~~ northeast ~~or to~~ southeast of the station, especially in summer, ~~but a~~
525 large area of sources was identified from the southwest to the northwest regions, which
526 also contained CH₄ emissions from the northwest Gansu province, the northwest

527 Qinghai province and the southeast of the XUAR. Additionally, the source regions
528 differed in various seasons as well as years (Fig. 6). The regions of potential source
529 regions in the spring (Fig. 6a-c) and winter (Fig. 6j-l) were obviously
530 larger than those in the summer (Fig. 6d-f) and autumn (Fig. 6g-i). The
531 seasonal difference was due to the effect of westerlies or the southeast monsoons
532 (Zhou et al., 2004).

533 There were also trends ~~infor~~ the CH₄ source regions ~~correlated~~ along with specific
534 years: i) the area of the potential source regions was increasing with the year over time,
535 and ii) the location of strong ~~potential~~ sources changed along with the time. For
536 example, in autumn and winter, the ~~strength of~~ CH₄ sources ~~were were~~ very strong in
537 the southeast ~~and to~~ northeast regions during 2004-2007 (Fig. 6g & i), and then
538 weakened in 2008-2012 (Fig. 6h & k). Finally, from and finally (i.e. 2013-2017
539 the sources), almost vanished in the eastern regions but moved to the southwest with a
540 very large distribution area (Fig. 6i & l).

541 More CH₄ sources appeared at the WLG along with the progression of time, which
542 could have been attributed to the influence of human expansion. The pattern of strong
543 sources moving indicated that the southwest area relative to the WLG, e.g. Northern
544 India, was gradually becoming a strong CH₄ contributor. India with abundant cattle as
545 well as an extensive large-scale coal mining operation possibly contributed large
546 amounts of CH₄ to move from northern India to the northeastern Tibetan Plateau (Fig.
547 7i & l) (Fu et al, 2012). The analysis of air mass transport (Fig. 6d) also supported the
548 conclusions that the air masses from the southwest regions contributed the highest CH₄
549 mole fractions. The studies of atmospheric Hg at the WLG by Fu et al. (2012) also
550 supported this phenomenon, which found the long-range transport of atmospheric Hg
551 from India to the Northeastern Tibetan Plateau.

553 **2.7 Extracting the well-mixed ambient methane**

554 ~~To precisely understand characteristics of atmospheric CH₄, e.g., seasonal cycle or~~

555 ~~long-term trend, it is vital to identify the CH₄ records that were influenced by local~~
556 ~~sources and sinks. In this study, we analyzed hourly CH₄ measurements during 1994–~~
557 ~~2017, and 47.3% of CH₄ data were classified as regional representative, with the~~
558 ~~average CH₄ mole fraction of 1847.9 ± 0.3 ppb. The local representative data was~~
559 ~~obviously larger than regional events, with an average value of 1858.2 ± 0.4 ppb (Table~~
560 ~~2). The proportion of regional events increased slightly before 2012, but significantly~~
561 ~~reduced in recent years (e.g., 2013–2017). The filtered regional/local time series was~~
562 ~~shown in Figure 7. It can be seen that the CH₄ mole fractions obviously increased from~~
563 ~~1994 to 2017. The atmospheric CH₄ showed strong growth and displayed large~~
564 ~~fluctuation at WLG (Fig. 7). In 1995, the average CH₄ mole fraction was only 1805.8~~
565 ~~± 0.1 ppb, however, the average value increased 98 ppb by the year of 2016 (1903.8 ±~~
566 ~~0.1 ppb) (Table 3).~~

568 ~~2.83.5~~ Correlation analysis between CH₄ and CO

569 Because parts of the CH₄ and CO in the atmosphere ~~are~~ were from the same
570 anthropogenic sources (e.g., e.g. fossil fuel combustion), the long-term trend of
571 ΔCO/ΔCH₄ was helpful to understand the variation in the sources/sinks in many studies
572 (Buchholz et al., 2016; Niwa et al., 2014; Tohjima et al., 2014; Wada et al., 2011). In
573 this study, the hourly CO data from 2004 to 2017 was used to further analyze the long-
574 term variations of CH₄ (Fig. S3). ~~Two~~ we calculated the regression slopes of ΔCO/ΔCH₄
575 from 2004 to 2017 ~~were analyzed~~ (Fig. S4) ~~Fig. S4~~. Figure ~~8-8~~ 8-8 presents the average
576 seasonal cycles of the ΔCO/ΔCH₄ slopes. Generally, the slopes were larger in
577 summer and small ~~lower~~ in winter during the observing period, except for 2004–2007
578 wherewith the highest slope was in autumn. This was primarily due to the effect of
579 monsoons and air mass transport. Tohjima et al. (2014) found an opposite variation at
580 the Hateruma Island, which showed small slope values in the summer. Wada et al.
581 (2011) analyzed more than 10-year seasonal variation of the ΔCO/ΔCH₄ ratios at three
582 monitoring stations, i.e. MNM, Yonagunijima (YON), and Ryori (RYO) in Japan,

583 which also showed an opposite trend to that of the WLG. It was because these sites
584 were considerably affected by the Asian continental source regions, where had
585 enhanced emissions of CH₄ in the summer (rice paddies) and CO in the winter (fuels
586 combustion). At the WLG, Regional polarization in the concentration ratios of CH₄ and
587 CO were seen (Fig. S5), which implied different strong source distributions between
588 CH₄ and CO. This result was also seen in the CPF analysis (Fig. 5). Furthermore, the
589 cluster results (Fig. 6b & d) and the PSCF analysis (Fig. 7f & l) also supported this
590 conclusion. In the summer, the source emissions were mainly from the east-southeast
591 regions (cities) with large amounts of CO but relatively lower CH₄ (Table 3), leading to
592 the largest $\Delta\text{CO}/\Delta\text{CH}_4$. In contrast, the sources mostly from southwest-west regions
593 (Tibetan Plateau) emitted large amounts of CH₄ but with relatively lower CO in the
594 winter. Hence, the opposite two air mass transports and source emissions led to a peak
595 in $\Delta\text{CO}/\Delta\text{CH}_4$ in the summer and a trough in the winter (Fig. 8). These results revealed
596 different local sources/sinks, the special topography conditions and source distributions
597 around the WLG.

598 ~~Additionally, the regression slopes increased along with the time, which showed~~
599 ~~the maximum in 2013-2017 and the minimum in 2004-2007. For the year to year~~
600 ~~variations, the $\Delta\text{CO}/\Delta\text{CH}_4$ slopes showed large fluctuations from 2004 to 2017 at the~~
601 ~~WLG (Fig. 9). The slopes showed a decreasing trend during 2004-2007 but then~~
602 ~~increased from 2007-2010, and again decreased again after 2010. In spring and~~
603 ~~summer, increasing trends appeared again after 2014. The slopes in summer were~~
604 ~~almost the largest but the lowest in winter.~~

605 In 2007, a large increase in ΔCH_4 appeared, and from 2010 to 2013, the ΔCO
606 decreased significantly (Fig. S4). Before 2010, large air masses and potential source
607 regions were identified in eastern regions (cities) with the highest CO emission (Fig. 7
608 and Table 3). After 2010, the southwest regions had the highest CH₄ emissions but
609 relatively low CO emissions. Therefore, the strong seasonal variation of $\Delta\text{CO}/\Delta\text{CH}_4$
610 also revealed that the WLG was affected by different anthropogenic sources, e.g.

611 sources from cities, and sources from the Tibetan Plateau, during a year, especially in
612 the summer and winter. The long-term trend of the slopes implied that the source
613 emission types (CO sources or CH₄ sources) around the WLG might have been
614 changing with human activities, like straw burning, in the early years, or coal mining
615 in recent years.

618 2.9.3.6 Variation of long-term records

619 2.9.13.6.1 Seasonal cycles

620 ~~In order to further investigate the characteristics of atmospheric CH₄, we divided the~~
621 ~~CH₄ observations into two main regions according to the above analysis, including~~
622 ~~geographical conditions, the effect of surface winds, the long-range transports and the~~
623 ~~potential source distributions. The first region was covered the northeast to southeast~~
624 ~~(NNE ... SE) of WLG, which was denoted as City Regions (CR). The second region~~
625 ~~was located south to west (S ... W) of the station and was well known Tibet (Qinghai~~
626 ~~Xizang) Plateau (TP) (Fig. S2). Accordingly, the hourly CH₄ records when the surface~~
627 ~~wind coming from these sectors were divided into two subsets (i.e. TP and CR). The~~
628 ~~long-term variations between the two regions as well as the total regional time series~~
629 ~~(Total) were compared and analyzed to explore new sight of atmospheric CH₄ variation~~
630 ~~at WLG.~~

631 Overall, ~~at WLG,~~ the seasonal averages of regionally representative CH₄ at the WLG
632 CH₄ were almost in a cycle starting with ordered by summer (18501861.7-0 ± 0.43 ppb),
633 then winter (185647.74 ± 0.3 ppb), autumn (185544.94 ± 0.43 ppb), and spring
634 (185344.12 ± 0.3 ppb), except during 1994-1997 where there was a with the maximum
635 in the winter and a minimum in the autumn (Fig. 10 Fig. 10). Seasonal averages in the
636 CR were slightly different from that in the TP and even the total regional data. The
637 seasonal averages in the TP were mostly higher than that in the CR from 1994 to 2019,

638 except for in the winter (Table S2). The maximum mole fractions were mostly found in
639 August and the minimum mole fraction appeared in April for the total regional time
640 series, with a seasonal amplitude of 14.4 ppb. The peak to trough amplitude in the CR
641 (~16.7 ppb) was higher than that in the TP (~15.1 ppb) during 1994-2019. The seasonal
642 variations were consistent with the previous short-term studies from 2002-2006 at
643 WLG (Zhang et al., 2013). However, these variations were almost opposite to
644 observations at the adjacent stations of Lin'an, Shangdianzi, and Longfengshan in
645 China (Fang et al., 2013, 2016). For other regional sites in the Asia, Guha et al. (2018)
646 studied seasonal variability at the Sinhagad (SNG) and Cape Rama station (CRI) over
647 India, which also showed an opposite trend to the WLG due to the strong impact of
648 monsoon dynamics. Ahmed et al. (2015) found that the seasonal CH₄ showed a
649 maximum in the winter and a minimum in the spring at two urban sites of Guro (GR)
650 and Nowon (NW), in Seoul, Korea over 2004-2013. Kim et al. (2015) investigated the
651 decadal variation (1991-2013) of CH₄ at the East Asian sites, e.g. Ulaan Uul (UUM) in
652 Mongolia and Tae-ahn Peninsula (TAP) in Korea, which revealed again an opposite
653 seasonal trend to that of the WLG.

654 The data were further compared with similar WMO/GAW global stations in the
655 Northern Hemisphere, including MLO (19.54° N, -155.58° E, 3397m a.s.l.)
656 (Dlugokencky et al., 1995, 2019a), JFJ (46.55° N, 7.99° E, 3580m a.s.l.) (Zellweger et
657 al., 2016), MNM (24.29° N, 153.98° E, 7.1m a.s.l.) (Matsueda et al., 2004; Tsutsumi et
658 al., 2006), as well as the marine boundary layer (MBL) from the NOAA/ESRL lab at a
659 similar latitude (Dlugokencky et al., 2019b). It was seen that the stations in the Northern
660 Hemisphere and the MBL showed an opposite trend to the WLG with the minimum
661 being in summer and maximum in the winter or spring (Fig. 11). The peak in the
662 summer at the WLG was attributed to larger grazing, human activities, ruminants, and
663 easterly winds coming from neighboring areas over other seasons. The CH₄ emissions
664 from yaks and other ruminants in the Tibetan Plateau (alpine pasture) were very strong
665 in the summer, preceded only by paddy emissions (Fang et al., 2013; Zhang et al.,

666 2013). Furthermore, the dynamic transport by airflow from the polluted
667 northeast/southeast region was also strong in summer, which all induced high CH₄ mole
668 fractions and consequently an opposite trend than other sites (Ma et al., 2002; Xiong et
669 al., 2009).

670 The seasonal amplitude at the WLG (~14 ppb) was significantly lower than many
671 other sites in the Northern Hemisphere, by about 35-70 ppb. Such sites included MLO
672 in America, BRW in North Pole, UUM in Mongolia, TAP in Korea, Ny-Ålesund in
673 Norway, Bialystok in Poland, Ochsenkopf in Germany, and Beromunster in Switzerland
674 (Dlugokencky et al., 1995; Kim et al., 2015; Morimoto et al., 2017; Thompson et al.,
675 2009; Popa et al., 2010; Satar et al., 2016). MBL also showed a larger amplitude than
676 WLG (Fig. 11). The study at the SNG and CRI over India showed a much larger
677 amplitude close to 200 ppb (Guha et al., 2018). The low amplitude at the WLG was
678 because of the high elevation of the continental mountain sites, where there were
679 relatively fewer effects from local influences than the coastal/island sites (Yuan et al.,
680 2019). Additionally, the Seasonal averages in CR were significantly different to that in
681 TP and also the entire regional data (Total). The seasonal average in TP was mostly
682 higher than that in CR, except for wintertime (Table S2). The atmospheric CH₄ in
683 August was mostly the maximum and the April was the minimum for the total regional
684 time series (Total), with the seasonal amplitude of 13.4 ppb. The peak to trough
685 amplitude in CR (~15 ppb) was higher than that in TP (~13 ppb) during 1994-2016.
686 Additionally, seasonal amplitudes indicated different trends between the CR and TP.
687 For the CR, the seasonal amplitude decreased atwas firstly _dropped_ and then
688 increased overalong with time, which were was similar to the variation of the total
689 regional events (Total). But for the TP, the amplitude hadisplayed a continuously
690 increasing trend, with values of about 15.19, 189.13, 21.46, 203.4, and 232.74 ppb in
691 1994-1997, 1998-2002, 2003-2007, 2008-2012, and 2013-20162019, respectively.

692 . This revealed that the Tibetan Plateau was intensively affected by strong regional
693 sources (e.g. grazing or emissions from India) over time.

694

695 2.9.23.6.2 Long-term trend

696 The fluctuating trend in atmospheric CH₄ during 1994-2019 at the WLG (Fig. 12) was
697 similar to the global trend reported by many studies (Bergamaschi et al., 2013; Rigby
698 et al., 2017; Nisbet et al., 2019). In the 1990s, the CH₄ growth rates were very low-small
699 and/or even negative at WLG. Subsequently, during 2002-2006, a steady period was
700 found with a near-zero growth rates was found during 2002-2006. However, After 2007,
701 the atmospheric CH₄ increaswas raiseded significantly after 2007 (Fig. 12Fig. 11a). In
702 the year of 1997/1998, 2000/2001, 2007/2008, and 2011/2012, a larger amplitude
703 fluctuation inof the growth rates was found and a strong growth appeared (Fig. 12Fig.
704 11b). The growth rates fluctuated evenly with both positive and negative values before
705 2009. However,But almost all of the growth rates showed a positive value after 2009.
706 From 1990s to 2010s, tThree apparently developing stages (i.e.i.e. highlighted green,
707 blue, and red blocks) were presentedcould be seen from the 1990s to 2010s, Tthe CH₄
708 mole fraction slightly decreased duringin 1998-2000 (green color), and then go-went
709 through a relatively steady period duringin 2003-2006 (blue color), finally increased
710 steadily-rapidly after 2007 (red color) (Fig. 12Fig. 11).

711 The ~~overall~~ annual growth rates ~~wasere~~ 5.1 ± 0.1 ppb yr⁻¹ ~~throughoutover~~ 1994-
712 ~~2016~~2019 at the WLG (Table 4Table 4). However, ~~t~~The periodic annual growth rates
713 were 4.96 ± 0.1 , 2.56 ± 0.2 , $4.95.3 \pm 0.12$, $7.76 \pm 0.1.2$ and 5.57 ± 0.1 ppb yr⁻¹ ~~duringin~~
714 1994-1997, 1998-2002, 2003-2007, 2008-2012, and 2013-~~2016~~2019, respectively.
715 Similar growth rates were found between the CR and the TP during 1994-2019 (Fig.
716 S6). In 1994-1997 and 2003-2007, the growth rates in the TP were even larger than that
717 in the CR (Table 4). These results indicated that there were also strong CH₄ sources
718 from the TP. The previous study by Zhou et al. (2004) showed the annual increase in
719 CH₄ by 4.5 ppb yr⁻¹ in 1992-2001, which was close to our study in 1994-1997 and 1998-
720 2002. Tohjima et al. (2002) found that the CH₄ levels at the Cape Ochi-ishi and
721 Hateruma Island in1995-2000 respectively increased by 4.5 and 4.7 ppb yr⁻¹, which

722 were also similar to that of the WLG. Tsutsumi et al. (2006) analyzed the trend of hourly
723 CH₄ data from 1998 to 2004 on the YON, which showed a similar increase (~3.0 ppb
724 yr⁻¹) to the WLG. The study at the GR and NW in Seoul, Korea, presented almost an
725 identical trend of 2 ppb yr⁻¹ between 2004 and 2013 (Ahmed et al., 2015), which was
726 lower than that of the WLG in similar period. In the early 1990s, the CH₄ growth rates
727 at the WLG were very low and similar to the global level. The level of •OH radicals
728 might control the decrease or increase of CH₄ in the atmosphere during this period
729 (Dlugokencky et al., 1998; Rigby et al., 2017; Turner et al., 2017). However, the growth
730 rates were high in 1998 (Fig. 12b), which may have been due to the high temperatures
731 and a large amount of biomass burning (Cunnold et al., 2002; Lelieveld et al., 2004;
732 Simmonds et al., 2005). The growth rate of CH₄ observed at the Ny-Ålesund, Svalbard,
733 increased from 0.3 ± 0.2 ppb yr⁻¹ during 2000-2005 to 5.5 ± 0.2 ppb yr⁻¹ during 2005-
734 2014, which had a similar variation but with a little lower growth rates than that of the
735 WLG (Morimoto et al., 2017). The study suggested that the temporal pause in 2000-
736 2005 was ascribed to the reductions of CH₄ emissions from the microbial and fossil fuel
737 sectors, while the increase in 2005-2014 was due to an increase in microbial release.

738 The constantly larger CH₄ growth rate after 2007 at the WLG (Fig. 12) (Table 4)
739 was similar to the recent studies by Nisbet et al. (2016) and (2019), which showed that
740 the global CH₄ increased by 5.7 ± 1.2 ppb yr⁻¹ in 2007-2014, and was much higher at
741 12.7 ± 0.5 ppb yr⁻¹ in 2014, with 10.1 ± 0.7 ppb yr⁻¹ in 2015, 7.0 ± 0.7 ppb yr⁻¹ in 2016,
742 and 7.7 ± 0.7 ppb yr⁻¹ in 2017. The average growth rate in the Northern Hemisphere
743 was 7.3 ± 1.3 ppb in 2007 and 8.1 ± 1.6 ppb in 2008 (Dlugokencky et al., 2009), which
744 was also similar to the observation at the WLG (Table 4). After 2007, most sites in the
745 Northern Hemisphere had large CH₄ growth rates. Also, the average global growth rate
746 was similar to the WLG at 7.1 ppb yr⁻¹ in the most recent ten years (WMO, 2019). Fang
747 et al. (2013) showed that the annual growth rate of CH₄ was 9.4 ± 0.2 ppb yr⁻¹ in 2009-
748 2011 at the WLG, which was a little higher than this study in 2008-2012. The adjacent
749 stations in China also revealed the high CH₄ growth rates of 8.0 ± 1.2 ppb yr⁻¹ at Lin'an

750 in 2009-2011, 7.9 ± 0.9 ppb yr⁻¹ at Longfengshan in 2009-2011, and 10 ± 0.1 ppb yr⁻¹
751 at Shangdianzi in 2009-2013 (Fang et al., 2013, 2016), which were all higher than the
752 similar periods of 2008-2012 or 2013-2019 at the WLG (Table 4). The CH₄
753 measurements in other countries, such as the Beromünster tall tower station, also
754 showed a high growth rate of 9.66 ppb yr⁻¹ in 2012-2014 (Satar et al., 2016). The warm
755 temperatures, biomass burning, and the climatic anomalies (El Niño or La Niña), likely
756 enhanced the CH₄ emissions after 2007 (Dlugokencky et al., 2009). The anomalous
757 years of increasing or decreasing (e.g. 2007/2008) might have a significant influence
758 on the overall CH₄ trend (Fig. 12). These frequent anomalies also appeared in most
759 long-term observation stations, e.g. MLO in the USA (Dlugokencky et al., 2009) and
760 Mount Zugspitze in Germany (Yuan et al., 2019), due to climatic forces, such as those
761 exceptions during the El Niño oscillation, forest fires, volcanic eruptions, and extreme
762 weather events (Keeling et al., 1995; Dlugokencky et al., 2009; Keenan et al., 2016;
763 Nisbet et al., 2019).

764 Many studies have investigated the potential reasons for the anomalous increasing.
765 The study by Satar et al. (2016) at Beromünster, Switzerland explained that the short-
766 term spikes were possibly related to emissions from agricultural activities, while the
767 longer-lasting peaks were because of air mass transport and mixing. The isotopic
768 evidence suggested that the significant increase of biogenic emissions was the dominant
769 factor for the CH₄ rise. This was especially true in the tropical wetlands that had strong
770 rainfall anomalies, and agricultural sources such as rice paddies and ruminants were a
771 cause while fossil fuel emissions were not the main cause (Nisbet et al., 2016). The
772 study from thaw ponds at Arctic regions revealed that there had very weak correlation
773 between the amount of CH₄ released from ponds and environmental factors, e.g. air
774 temperature and atmospheric pressure (Burke et al., 2019). Sweeney et al. (2016) using
775 29-year of measurements on North Slope of Alaska (BRW) to investigate the sensitivity
776 of CH₄ emissions to the temperature change, which revealed that despite the short-term
777 temperature sensitivity increases CH₄ emissions, it would have little impact in the long

778 term. However, up to now, the specific causes of such distinct variability through the
779 years, including the spikes or near-zero CH₄ growth rates, have not yet been determined.
780 It is well established that human activities were mainly responsible for the recent rapid
781 CH₄ growth rates and anomalies. The analysis from EDGAR showed the CH₄ emission
782 per sector in China (Fig. S7) (Crippa et al., 2019). During the observing period, the
783 waste, oil, and natural gas combustion, and open burning continuously emitted large
784 amounts of CH₄ into the air. After 2000, the CH₄ emissions from solid fuel increased
785 greatly in China. After 2003, the CH₄ emitted from rice cultivation also increased
786 continuously (Fig. S7). The increased emissions from these sectors greatly contributed
787 to the CH₄ increase at the WLG, as well as the other regions in China. In addition,
788 studies revealed that China's coal sector dominated the positive trend in recent years,
789 which contributed to the highest proportion of anthropogenic CH₄ emissions (~33%)
790 (Janssens-Maenhout et al., 2019; Miller et al., 2019). In 2010-2015, China's coal
791 production increased (from 3400 to 4000 million metric tons), but CH₄ emissions from
792 rice cultivation, agriculture practices, ruminants, waste, and oil/gas consumption only
793 increased slightly if at all (EIA, USA). Therefore, the control measures of coal mining
794 reduction or limiting natural gas and petroleum exploitation may play an important role
795 in slowing down CH₄ emissions in China.

796 ~~The CH₄ growth rate in CR was significantly different from that in TP (Fig. S3).~~
797 ~~In 1994-1997, 2003-2007 and 2013-2016, the growth rates in TP were obviously larger~~
798 ~~than that in CR (Table 4). But in 2003-2007 and 2008-2012, the CR showed higher~~
799 ~~annual growth rates. In addition, for the entire observing period (i.e. 1994-2016), the~~
800 ~~growth rates in both TP (5.2 ± 0.1 ppb yr⁻¹) and CR (5.0 ± 0.1 ppb yr⁻¹) were similar to~~
801 ~~the overall annual growth rates (Fig. S3).~~

803 2.103.7 Case study for air mass transport

804 As described above, the northeast and southeast city regions might have acted as strong
805 regional sources influencing the atmospheric CH₄ at the WLG. Therefore, to analyze

806 the effect of long-~~distance-range~~ transport of emissions from cities, ~~we further excluded~~
807 the regionally ~~representative data~~ ~~by air mass transport~~ ~~was further excluded by air mass~~
808 ~~transport~~, and ~~the the rest of remaining~~ regional records ~~were-were~~ denoted as ‘TR’.
809 ~~First, We applied~~ the monthly cluster analysis ~~was applied~~ to hourly trajectories over
810 2005-2007, 2008-2012 and 2013-2017. ~~Then, based on the cluster analysis, the clusters~~
811 ~~were divided into two groups, i.e. from city regions (red clusters in Fig. S8), and other~~
812 ~~(black clusters in Fig. S8). Finally, the regionally representative data were accordingly~~
813 ~~classified as two groups based on the cluster results (cities or other).~~ The ~~statistical~~
814 ~~result~~ ~~cluster results from city regions~~ were ~~presented-presented~~ in detail in ~~Fig.~~
815 ~~S4~~ ~~Figure S8~~ and ~~Table S3~~ ~~Table S3~~.

816 ~~Consequently, t~~The proportions of trajectories from cities ~~were~~ ~~as~~ 40.3%, 32.5%,
817 and 6.8% in 2005-2007, 2008-2012, and 2013-2017, respectively. ~~And~~ ~~about~~ 22.8%,
818 35.6%, and 38.4% of the regional records were associated with air masses ~~transport~~
819 from city regions in 2005-2007, 2008-2012, and 2013-2017, respectively (~~Table 5~~ ~~Table~~
820 ~~5~~). The average CH₄ values ~~for the when air masses transport from~~ city region ~~air~~
821 ~~masses~~ (1863.0 ± 0.3 ppb) ~~were~~ ~~obviously~~ higher than ~~the~~ other sectors ($1850.6 \pm$
822 0.2 ppb) (~~Fig. S9~~ ~~Fig. S5~~). ~~The overall growth rates of the TR in the periods of 2005-~~
823 ~~2016 or 1994-2016 were similar to the original data series (Fig. S9).~~ However, after
824 excluding the CH₄ records ~~for the when~~ air trajectory~~ies~~ ~~s~~ transports from city regions,
825 the growth rates of ~~the~~ TR in 2008-2012 (10.1 ± 0.1 ppb yr⁻¹) and 2013-2017 (6.3 ± 0.1
826 ppb yr⁻¹) (~~Table 5~~ ~~Table 5~~) were higher than the original regional data series (~~i.e.i.e.~~ 7.6
827 ± 0.2 and 5.7 ± 0.1 ppb yr⁻¹) (~~Table 4~~ ~~Table 4~~). ~~And the overall growth rate for TR in~~
828 ~~2005-2016 or 1994-2016 was still similar to original data series (Total), no significant~~
829 ~~difference was found (Fig. S5).~~

830 ~~These results suggested that there were possibly other strong CH₄ sources at the~~
831 ~~WLG that were not from cities and the southwest region (Northern India) was the most~~
832 ~~likely contributor. The PSCF analysis also supported this result (Fig. .7). At present,~~
833 ~~Northern India and Eastern China were the two largest sources of CH₄ at the WLG (Fig.~~

834 S10) (Crippa et al., 2019). Since the Tibetan Plateau was coincidentally trapped in the
835 middle of them, the atmospheric CH₄ at WLG was very likely dominated by long-range
836 transport from these two regions. Although CH₄ emissions increased slowly during
837 1994-2002, a negative trend appeared (Fig. S10), significantly increased emissions
838 were found in both southeast and southwest Asia after 2007. Chen et al. (2013)
839 illustrated that the warming (0.2 °C per decade) in the Tibetan Plateau resulted in
840 substantial emissions of CH₄ due to the thawed permafrost and melted glaciers. The
841 rapid increase of CH₄ would probably make it difficult to meet the goals of carbon
842 emission reduction in the future. This would be especially true with the scenario of
843 quickly increasing CH₄ on the Qinghai-Tibetan Plateau due to the emissions from the
844 two largest source regions of Northern India and Eastern China. The large growth rate
845 of atmospheric CH₄ in the TP revealed that i) the atmospheric CH₄ at the WLG was not
846 predominantly influenced by eastern cities in recent years and ii) large amounts of CH₄
847 were transported from the Tibetan Plateau to WLG in recent years.

848

849 ~~–Discussion~~

850 ~~–Anthropogenic emission on temporal patterns~~

851 ~~In the early years, the daily cycles of atmospheric CH₄ at WLG were not distinct~~
852 ~~and the amplitude was small (Fig. 2a b), which were similar to the previous studies by~~
853 ~~Zhang et al. (2013) and Fang et al. (2013). The ambiguous diurnal patterns indicated~~
854 ~~that the local sources were weak at WLG in the past. The apparent diurnal cycles after~~
855 ~~2000 may be attributed to the intense activities by human (e.g., grazing, fuel burning),~~
856 ~~which was aggravated in daytime and weak in nighttime (Fang et al., 2013). The~~
857 ~~meteorological conditions (e.g., diffusion and transport) could contribute to the~~
858 ~~increasing CH₄ amplitudes. The WLG was remote from the populate center, the good~~
859 ~~diffusion condition in daytime may bring high CH₄ mole fractions to the site. The~~
860 ~~increasing amplitude (Table S1) and CH₄ mole fractions over time suggested that the~~
861 ~~WLG was affected by increasingly local sources (e.g., human activities) (Zhou et al.,~~
862 ~~2004). The maximum was found in summer in recent years (Fig. 2e) could also be~~
863 ~~ascribed to the transport of anthropogenic sources by the meteorological factors. In~~
864 ~~summer, the intensely herd or graze activities around WLG might enhance the regional~~
865 ~~CH₄ emissions and hence contribute to the higher CH₄ mole fractions. The higher CH₄~~
866 ~~mole fraction in winter in past years (Fig. 2a) was probably because of the way of~~
867 ~~heating (e.g., large biomass burning) as well as the adverse diffusion conditions in cold~~
868 ~~weather.—~~

869 The previous study by Zhou et al. (2004) found that higher CH₄ mole fractions
870 appeared when the winds come from the ENE-E-ESE-SE sectors at WLG during 1991-
871 2002, which was similar to our study in similar period (Fig. 3a b). Causes of the
872 elevated CH₄ from these wind sectors could be attributed to the large plantation of
873 highland barley as well as high population density in those areas (Fang et al., 2013).
874 Two largest cities Xining and Lanzhou are situated in northeast and east of WLG,
875 respectively, which could emit large amount of CH₄ by human activities. Besides,
876 previous studies on black carbon (BC) and carbon monoxide (CO) indicated that the
877 emissions from the Yellow River Canyon industrial area, ~500km away from northeast
878 of WLG may also donate to the high CH₄ values originating from ENE and NE sectors
879 (Zhou et al., 2003). In summer, the prevailing wind directions were from NE... ESE
880 sectors (~46%) (Fig. S6), and the CH₄ mole fractions were also higher in the related
881 sectors. However, in the autumn and winter, although the prevailing wind and high wind
882 speed were from SSW... W sectors (~40-50%) (Fig. S6), the high CH₄ mole fractions
883 were from almost the opposite wind sectors of NNE... ESE (Fig. 3f), which indicated
884 that strong local sources were distributed from northeast to southeast (city regions), and
885 even covered the emissions of natural sources. As time goes on, the wind sectors with
886 high CH₄ mole fractions changed and concentrated on ESE to ENE sectors, and the
887 amplitudes of enhancements were increasing, which further implied the effect of
888 stronger emissions from anthropogenic sources in city regions in recent years.

889 ~~Pollutant sources regions~~

890 ~~Sources regions~~

891
892 The air masses from east and northeast regions passed over the cities of Xining
893 and Lanzhou (capital of Gansu province), which is the populated center and industrial
894 area (Fig. 5). However, the highest CH₄ values was not observed when air mass was
895 from these sectors. Instead, high CH₄ mole fractions were frequently observed when air
896 mass from the northwest to southwest (Table 1). It was possibly due to that the air
897 masses from west and northwest had passed through the northwest of Qinghai province
898 and the central area of Xinjiang Uygur Autonomous Region (XUAR), where located
899 Ge'ermu urban area (the second largest city of Qinghai) with rapid industrial
900 development, natural gas and petroleum resources exploitation and large crops residue
901 burning (Zhang et al., 2013). Similar to the CPF percentile analysis (Fig. 4), the
902 southwest or northwest region away from the site may be also strong source regions.

903 Most potential source identified in northwestern regions (Fig. 6) was possibly due
904 to CH₄ emissions from the northwest Gansu province, the northwest Qinghai province
905 and the southeast of XUAR. The different source distribution by seasons could be
906 attributed to the effect of westerlies or the southeast monsoons (Zhou et al., 2004). The
907 obviously increasing source region was clear evidence for the strong effect of the
908 expansion of human activity. Moreover, the pattern of source region moved from the
909 east to the southwest, especially in autumn and winter, indicated that the southwest
910 away from the WLG, e.g., Northern India, were gradually becoming a strong CH₄

911 source region. India has abundant cattle as well as extensive large scale coal mining,
912 large amount of CH₄ emissions may transport from northern India to the northeastern
913 Tibetan Plateau (Fig. 6i & l). The air mass transport result (Fig. 5d) also support the
914 result that the southwest air masses (cluster 1) contributed the highest CH₄ mole
915 fractions. The studies of atmospheric Hg at WLG by Fu et al. (2012) also revealed this
916 phenomenon.—

917 —Different sources between CH₄ and CO

918 The percentile polar plot clearly showed the specific distribution of different CH₄
919 mole fractions (Fig. 4). The result revealed that most areas around WLG contributed to
920 low CH₄ mole fractions, the southeast and southwest of the site exist two strong source
921 regions. It is of great possible that the anthropogenic emission from cities (e.g.,
922 Lanzhou, Chengdu, etc.) was the only cause for high values in the southeast, and the
923 southwest region away from WLG was possibly due to sources from other countries,
924 such as India. Unlikely to the CH₄, the high CO mole fractions were consistent from
925 east regions (urbanized area) (Fig. 4), indicating strong anthropogenic sources in city
926 regions (i.e. Xining and Lanzhou) (Zhang et al., 2011).—

927 The seasonal cycles of $\Delta\text{CO}/\Delta\text{CH}_4$ slopes (high in summer and low in winter) (Fig.
928 8) may primarily be due to the effect of monsoons and air mass transport. In summer,
929 air masses arriving at WLG were predominantly transported from the northeast to east
930 city regions (e.g., Xining, Lanzhou) with the largest CO mole fractions. In contrast, the
931 air masses were mainly from the southwest in winter, which carried strong CH₄
932 emissions but few CO emissions (Zhang et al., 2011) (Table 1). Hence the opposite two
933 air mass transport lead to a peak in summer and a trough in winter (Fig. 8). Moreover,
934 we could see apparently regional polarization in the concentration ratio of CH₄ and CO
935 (Fig. S7), implying the different strong source distribution between CH₄ and CO at
936 WLG. The cluster results (Fig. 5b & d) and the potential sources analysis (Fig. 6f & l)
937 also support this seasonal variation. Tohjima et al. (2014) found an opposite variation
938 at Hateruma Island, which showed low slope values in summer. It could be attributed
939 to different local sources and sinks, suggested the special topography condition and
940 local source distribution around WLG. The large $\Delta\text{CO}/\Delta\text{CH}_4$ fluctuations (Fig. 9) over
941 the study period was likely because of the anomaly years of different CH₄ or CO mole
942 fractions as well as source regions. In 2007, large increase of ΔCH_4 appeared, and from
943 2010 to 2013, the ΔCO decreased significantly (Fig. S1). Before 2010, large air masses
944 and potential source regions were identified in eastern regions (city regions) with the
945 highest CO emission (Fig. 5; Fig. 6). After 2010, the southwest regions showed high
946 contributions, with the highest CH₄ emission but relatively lower CO emission.
947 Therefore, obviously variation of the slopes presented with almost an increase in 2007-
948 2010 and a decrease after 2010 (Fig. 9).—

949 —Long term variations—

950 —Seasonal cycles

953 The seasonal variations with maximum in summer and minimum in spring at WLG
954 (Fig. 10) were consistent with the previously short-term studies, e.g., Zhang et al.
955 (2013) in 2002–2006. However, it was almost opposite to observation in the adjacent
956 stations such as Lin'an, Shangdianzi and Longfengshan in China (Fang et al., 2013,
957 2016). We further compared similar WMO/GAW global stations in the north
958 hemisphere, including MLO (19.54° N, 155.58° E, 3397m a.s.l.) (Dlugokencky et al.,
959 1995, 2019a), JFJ (46.55° N, 7.99° E, 3580m a.s.l.) (Zellweger et al., 2016), MNM
960 (24.29° N, 153.98° E, 7.1m a.s.l.) (Matsueda et al., 2004; Tsutsumi et al., 2006), as well
961 as the marine boundary layer (MBL) from NOAA/ESRL lab at similar latitude
962 (Dlugokencky et al., 2019b). as well as most global stations in the northern hemisphere,
963 e.g., MLO, BRW and JFJ (Dlugokencky et al., 1995; Loov et al., 2008). We further
964 compared similar WMO/GAW global stations in the north hemisphere, including MLO
965 (19.54° N, 155.58° E, 3397m a.s.l.) (Dlugokencky et al., 1995, 2019a), JFJ (46.55° N,
966 7.99° E, 3580m a.s.l.) (Zellweger et al., 2016), MNM (24.29° N, 153.98° E, 7.1m a.s.l.)
967 (Matsueda et al., 2004; Tsutsumi et al., 2006), as well as the marine boundary layer
968 (MBL) from NOAA/ESRL lab at similar latitude (Dlugokencky et al., 2019b). It can be
969 seen that the stations in the northern hemisphere (i.e. MLO, JFJ and MNM) and the
970 MBL showed an opposite trend to WLG with the minimum in summer and maximum
971 in winter or spring (Fig. 12). And the seasonal amplitude at WLG (~14 ppb) was lower
972 than many other sites in the northern hemisphere, about 35–70 ppb, e.g., MLO, BRW,
973 bialystok in Poland, ochsenkopf in Germany and beromunster in switzerland
974 (Dlugokencky et al., 1995; Thompson et al., 2009; Popa et al., 2010; Satar et al., 2016).
975 MBL also showed larger amplitude than that in WLG (Fig. 12). The low amplitude at
976 WLG was possibly because of high elevation of continental mountain sites (e.g., WLG,
977 JFJ), where were relatively less affected by local influences than the coastal/island sites
978 (Yuan et al., 2019). The peak in summer at WLG may be attributed to larger grazing
979 and human activities than other seasons. The CH₄ emissions from yaks and other
980 ruminants in Tibetan Plateau (alpine pasture) were very strong in summer, preceded
981 only by paddy emission (Fang et al., 2013; Zhang et al., 2013). Furthermore, the
982 photochemical capacities were very weak (high altitude) and the dynamic transport by
983 air flow from polluted northeast/southeast region was also strong in summer at WLG,
984 which all induced high CH₄ mole fractions in summer and consequently an opposite
985 trend with other sites (Ma et al., 2002; Xiong et al., 2009).—

986 —Long-term trends in different observing periods—

987 The entirely fluctuant trend of atmospheric CH₄ in 1994–2016 at WLG (Fig. 11)
988 was similar to the global trend reported by quite a few studies (Bergamaschi et al., 2013;
989 Rigby et al., 2017; Nisbet et al., 2019). The previous study by Zhou et al. (2004) showed
990 the CH₄ annual increase of 4.5 ppb yr⁻¹ in 1992–2001, which was similar to that in 1994–
991 1997 (4.6 ± 0.1 ppb yr⁻¹) and 1997–2002 (2.6 ± 0.2 ppb yr⁻¹) in our study (Table 4).
992 Tohjima et al. (2002) found similar growth rates that the CH₄ at Cape Ochi ishi and
993 Hateruma Island in 1995–2000 was increased about 4.5 and 4.7 ppb yr⁻¹, respectively. In
994

995 early 1990s, the CH₄ trend at WLG is very low, which was similar to the global growth
996 rates (Fig. 11b). The levels of OH radicals may have controlled the decrease or increase
997 of CH₄ in the atmosphere during this period (Dlugokencky et al., 1998; Rigby et al.,
998 2017; Turner et al., 2017). The growth rates were high in 1998 (Fig. 11b), which may
999 have been due to the high temperatures, large biomass burning and weak destruction
1000 (Cunnold et al., 2002; Lelieveld et al., 2004; Simmonds et al., 2005).

1001 The continuously larger CH₄ growth rates after 2007 at WLG (Fig. 11), e.g., $7.6 \pm$
1002 0.2 ppb yr⁻¹ in 2008–2012, 5.7 ± 0.1 ppb yr⁻¹ in 2013–2016 (Table 4), were similar to the
1003 recent study by Nisbet et al. (2016) and (2019), which showed the global CH₄ increased
1004 by 5.7 ± 1.2 ppb yr⁻¹ in 2007–2014, and the much higher of 12.7 ± 0.5 ppb yr⁻¹ in 2014,
1005 10.1 ± 0.7 ppb yr⁻¹ in 2015, 7.0 ± 0.7 ppb yr⁻¹ in 2016 and 7.7 ± 0.7 in 2017 ppb yr⁻¹.
1006 The average growth rate in the northern hemisphere was 7.3 ± 1.3 ppb in 2007 and 8.1
1007 ± 1.6 ppb in 2008 (Dlugokencky et al., 2009), which was also similar to the observation
1008 at WLG (Table 4). After 2007, most sites in the northern hemisphere displayed large
1009 CH₄ growth rates. Fang et al. (2013) showed that the annual growth rate of CH₄ was
1010 9.4 ± 0.2 ppb yr⁻¹ in 2009–2011 at WLG, which were a little higher than our study in
1011 similar period. The adjacent stations in China also revealed high CH₄ growth rates, e.g.,
1012 8.0 ± 1.2 ppb yr⁻¹ at Lin'an in 2009–2011, 7.9 ± 0.9 ppb yr⁻¹ at Longfengshan in 2009–
1013 2011, and 10 ± 0.1 ppb yr⁻¹ at Shangdianzi in 2009–2013 (Fang et al., 2013, 2016),
1014 which was higher than the similar period in 2008–2012 or 2013–2016 at WLG (Table
1015 4). The CH₄ measurements in other countries, e.g., Beromünster tall tower station also
1016 showed high growth rate of 9.66 ppb yr⁻¹ in 2012–2014 (Satar et al., 2016). The very
1017 warm temperatures, large biomass burning and the climatic anomaly e.g., El Niño, La
1018 Niña event, were likely enhanced the CH₄ emissions after 2007 (Dlugokencky et al.,
1019 2009). Additionally, the anomalously sharply increasing or decreasing years (e.g.,
1020 2007/2008) may have a significant influence on the overall CH₄ trend (Fig. 11), and
1021 these frequent anomalies appeared in most long-term observation stations, e.g., MLO
1022 in USA (Dlugokencky et al., 2009), Mount Zugspitze in Germany (Yuan et al., 2019),
1023 which were also possibly attributed to climatic forces, such as the exception during the
1024 El Niño oscillation, forest fires, volcanic eruptions, and extreme weather events
1025 (Keeling et al., 1995; Dlugokencky et al., 2009; Keenan et al., 2016; Nisbet et al., 2019).
1026 The study by Satar et al. (2016) at Beromünster, Switzerland explained that the short-
1027 term spikes were possibly related to emissions from agricultural activities, while the
1028 longer-lasting peaks were because of air mass transport and mixing.

1029 It is well established that the human activities are mainly responsible for the recent
1030 rapid CH₄ growth rates and anomalies. The analysis from Emissions Database for
1031 Global Atmospheric Research (EDGAR) showed the CH₄ emission by per sector in
1032 China (Fig. S8) (Crippa et al., 2019). During the observing period, the waste, the oil
1033 and natural gas and open burning continuously emitted large amount of CH₄ into the
1034 air. After 2000, the CH₄ emission from solid fuel increased greatly in China. After 2003,
1035 the CH₄ emitted from rice cultivation also increased continuously (Fig. S8). The
1036 increased emissions from these sectors may greatly contribute to the CH₄ increase at

1037 WLG, as well as the other regions in China. In addition, the recent studies revealed that
1038 China's coal sector may have dominated the clearly positive trend in recent years, which
1039 contributed the highest proportion of the anthropogenic CH₄ emissions (~33%)
1040 (Janssens Maenhout et al., 2019; Miller et al., 2019). In 2010–2015, China's coal
1041 production increased obviously (from 3400 to 4000 million metric tons), but emissions
1042 trends of CH₄ by rice, agriculture, ruminants, waste, and oil/gas have grown slightly
1043 and even remained flat (EIA, USA). The isotopic evidence suggests that the significant
1044 increase of biogenic emissions was the dominant factor of CH₄ rise, especially in the
1045 tropical wetlands with strong rainfall anomalies, or the agricultural sources such as rice
1046 paddies and ruminants, fossil fuel emissions have not been the main cause (Nisbet et
1047 al., 2016). The study by Chen et al. (2013) illustrated that the warming (0.2 °C per
1048 decade) in the Tibetan Plateau resulted in substantial emission of CH₄ due to the
1049 permafrost thawing and glaciers melting. However, up to now, the specific causes of
1050 such distinct variability around the years, e.g., the spikes or near-zero CH₄ growth rates,
1051 are not yet determined.

1052 –Annual growth rate in Qinghai-Tibetan Plateau

1053 Although similar annual growth rates were found among the City Regions (CR), the
1054 Tibet Plateau (TP) and total regional records (Total) in the entire observing period
1055 (1994–2016) (Fig. 11), significant differences were found in short-term periods (Fig.
1056 S3). In 2013–2016 (Table 4), the TP showed larger growth rate than that in CR, implying
1057 stronger CH₄ source in the Tibetan Plateau in recent years. The seasonal amplitude in
1058 the Tibetan Plateau was continuously increasing, which also revealed that the Tibetan
1059 Plateau was intensively affected by strong regional sources. Without air mass transport
1060 from the city regions, the significantly increased annual growth rate (TR) in 2008–2012
1061 (10.1 ± 0.1 ppb yr⁻¹) and 2013–2016 (6.3 ± 0.1 ppb yr⁻¹) (Fig. S5 and Table 5) suggested
1062 that there were possibly other strong CH₄ sources around WLG not from cities.
1063 Northern India and eastern China were obviously the largest two source regions of CH₄
1064 at WLG (Fig. S9) (Crippa et al., 2019). Since the Tibetan Plateau was coincidentally
1065 trapped in the middle of the largest increased source areas, the atmospheric CH₄ at
1066 WLG was very likely dominated by long distance transport from these two regions.
1067 Although CH₄ emissions increased slowly during 1994–2002, and even negative trend
1068 appeared in southeast China (Fig. S9), significantly increased emissions appeared in
1069 both southeast and southwest Asia after 2007. The rapidly increased CH₄ would
1070

1071 probably make it difficult to meet the goals of carbon emission reduction in the future.
1072 Especially on the scenario of quick increasing CH₄ on the Qinghai-Tibetan Plateau due
1073 to the emission from two largest source regions. In view of the integrated eco-
1074 environmental change processes and unique topography in the Qinghai-Tibetan
1075 Plateau, it may provide us one of the last precious regions to study global climate
1076 changes (Chen et al., 2013). The anomalously year to year fluctuations of atmospheric
1077 CH₄ in Tibetan Plateau were unquestionably a warning or alarm to the world, and the
1078 unprecedented annual growth rate might be a dangerous signal to global climate change.

1079

1080 3.4 Conclusions

1081 Three developing stages of atmospheric CH₄ at Mt. Waliguan from the 1990s to 2010s
1082 were found. The CH₄ atmospheric CH₄ at Mt. Waliguan increased continuously during
1083 1994-2017. mole fractions slightly decreased during 1998-2000, and then went
1084 through a relatively steady period during 2003-2006 and finally increased rapidly after
1085 2007. Although near-zero and even negative growth appeared in some ~~particular~~
1086 periods, e.g., 1999-2000, and 2004-2006, the overall trend of CH₄ was increased
1087 rapidly, especially in recent ~~decade~~ years. Obvious diurnal cycle was found with the
1088 peak at noon and a trough at late afternoon. Due to the unique geophysical locations
1089 and transport pathway, the seasonal averages of CH₄ at WLG displayed an opposite
1090 trend with sites in the northern hemisphere, with summer maximum and spring
1091 minimum. Although most areas around the WLG had small contributions to the CH₄
1092 emissions, two strong source regions were found from the northeast and southwest of
1093 the site. Large amount of air masses was from west and northwest regions of WLG,
1094 which accompanied with higher CH₄ mole fractions than that from city regions. The
1095 Northern India possibly has become a stronger source of CH₄ contributor to WLG
1096 rather than city regions before were in the past.

1097 ~~As time goes by, the~~ temporal patterns (e.g., seasonal amplitude), the annual
1098 variations, the long-term trends, or the potential source distribution of CH₄ at WLG

1099 ~~were all changing~~changed in recent years. We found that the WLG was increasingly
1100 ~~affected by local sources such as human activities. Thus, the long-term verification is~~
1101 ~~extremely important to accurately understand CH₄ variations. Additionally, the Tibetan~~
1102 ~~Plateau was intensively affected by strong sources over time~~The case study in Qinghai-
1103 ~~Tibetan Plateau revealed unprecedented annual growth rates of CH₄. In recent years,~~
1104 ~~which~~the Tibetan Plateau ~~_ even~~ showed a larger growth rate than that ~~in of the~~ city
1105 regions ~~in some periods.~~ ~~_~~ ~~Tibetan Plateau was with the highest average altitude and~~
1106 ~~was almost impervious to strong human activities. There is no doubt that the~~
1107 ~~anomalously~~ variation and ~~the~~ unprecedented ~~annual~~ growth rate of ~~the~~ atmospheric
1108 CH₄ in this region ~~might be a dangerous signal to global climate change~~revealed that it
1109 ~~was urgent to control CH₄ emissions. Reducing the emissions from strong source~~
1110 ~~sectors like coal mining, natural gas or solid fuel exploitation, and rice cultivation may~~
1111 ~~play an important role on CH₄ emissions reduction in China.~~

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1117 *Data availability.* The gridded meteorological data (2004-2017) from NOAA-ARL was
1118 available at <ftp://arlftp.arlhq.noaa.gov/pub/archives/gdas1/>. The data from MLO, JFJ,
1119 MNM station was downloaded from World Data Centre for Greenhouse Gases
1120 (WDCGG) at <https://gaw.kishou.go.jp/>. The MBL data was available at
1121 [ftp://aftp.cmdl.noaa.gov/data/trace_gases/CH₄/flask/surface/](ftp://aftp.cmdl.noaa.gov/data/trace_gases/CH4/flask/surface/). The geographical
1122 distribution of annual emission data by Emissions Database for Global Atmospheric
1123 Research (EDGAR) was from website
1124 https://edgar.jrc.ec.europa.eu/overview.php?v=50_GHG.

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1126

1127 *Author contributions.* SL, SF and ZF designed the research. SL performed the data
1128 processing with assistance of SF and MG. The station were monitored, maintained ML
1129 and PL, and they collected, preprocessed, and provided the hourly observational
1130 dataset. SL and SF finished the manuscript with contributions from all the co-authors.

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1132

1133 *Competing interests.* The authors declare that they have no conflict of interest.

1134

1135

1136 *Acknowledgments.* This study was funded by the National Key Research and
1137 Development Program of China (2017YFC0209700). We also thanks to the staff who
1138 have contributed to the system installation and maintenance at the Waliguan in past
1139 decades.

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1480 **Table 1.** The statistics of the filtered CH₄ data series over different periods during 1994-2019 at the
 1481 WLG station.

<u>Year</u>	<u>Regionally representative</u>			<u>Locally influenced</u>		
	<u>Hours</u>	<u>Percentage (%)</u>	<u>Mean (ppb)</u>	<u>Hours</u>	<u>Percentage (%)</u>	<u>Mean (ppb)</u>
<u>1994-1997</u>	<u>16122</u>	<u>71.3</u>	<u>1801.7 ± 0.5</u>	<u>6481</u>	<u>28.7</u>	<u>1806.2 ± 0.3</u>
<u>1998-2002</u>	<u>26347</u>	<u>83.2</u>	<u>1832.6 ± 0.7</u>	<u>5336</u>	<u>16.8</u>	<u>1837.7 ± 0.3</u>
<u>2003-2007</u>	<u>28181</u>	<u>69.4</u>	<u>1832.3 ± 0.3</u>	<u>12443</u>	<u>30.6</u>	<u>1839.2 ± 0.2</u>
<u>2008-2012</u>	<u>19627</u>	<u>63.5</u>	<u>1856.2 ± 0.4</u>	<u>11287</u>	<u>36.5</u>	<u>1865.2 ± 0.3</u>
<u>2013-2019</u>	<u>21683</u>	<u>44.2</u>	<u>1906.8 ± 0.3</u>	<u>27329</u>	<u>55.8</u>	<u>1920.4 ± 0.4</u>
<u>1994-2019</u>	<u>111960</u>	<u>64.0</u>	<u>1865.8 ± 0.4</u>	<u>62876</u>	<u>36.0</u>	<u>1868.2 ± 0.3</u>

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Table 2. Yearly average CH₄ mole fractions at the WLG station.

Year	Mean (ppb)	year	Mean (ppb)
<u>1994</u>	<u>1799.0 ± 0.4</u>	<u>2007</u>	<u>1837.2 ± 0.5</u>
			<u>1854.8 ±</u>
1995	180 <u>35.68</u> ± 0.1	<u>2008</u> 2006	<u>0.14835.6 ±</u> <u>0.2</u>
			<u>1847.2 ±</u>
1996	180 <u>84.86</u> ± 0.2	<u>2009</u> 2007	<u>0.14839.9 ±</u> <u>0.5</u>
			<u>1856.6 ±</u>
1997	18 <u>1106.58</u> ± 0.2	<u>2010</u> 2008	<u>0.24865.9 ±</u> <u>0.3</u>
			<u>1867.4 ±</u>
1998	182 <u>67.80</u> ± 0.1	<u>2011</u> 2009	<u>0.14849.1 ±</u> <u>0.4</u>
			<u>1879.6 ±</u>
1999	18 <u>1920.72</u> ± 0.1	<u>2012</u> 2010	<u>0.24857.9 ±</u> <u>0.2</u>
			<u>1895.7 ±</u>
2000	18 <u>1920.70</u> ± 0.2	<u>2013</u> 2011	<u>0.44872.6 ±</u> <u>0.4</u>
			<u>1890.2 ±</u>
2001	184 <u>79.32</u> ± 0.4	<u>2014</u> 2012	<u>0.24881.2 ±</u> <u>0.3</u>
			<u>1913.0 ±</u>
2002	183 <u>35.75</u> ± 0.2	<u>2015</u> 2013	<u>0.44896.3 ±</u> <u>0.2</u>
			<u>1914.4 ±</u>
2003	184 <u>02.85</u> ± <u>0.23</u>	<u>2016</u> 2014	<u>0.24890.4 ±</u> <u>0.4</u>
			<u>1911.6 ±</u>
2004	1836. <u>17</u> ± <u>0.24</u>	<u>2017</u> 2015	<u>0.14905.6 ±</u> <u>0.3</u>
			<u>1925.6 ±</u>
2005	183 <u>67.74</u> ± 0.1	<u>2018</u> 2016	<u>0.34903.8 ±</u> <u>0.4</u>
<u>2006</u>	<u>1834.7 ± 0.2</u>	<u>2019</u>	<u>1932.0 ± 0.1</u>

1485 **Table 13.** The statistics for the cluster analysis result for both CH₄ and CO from 2004-2017 at the
 1486 WLG station. The clusters from urban areas are highlighted with facein bold.

	Cluster	Number	Average CH ₄ mole fraction	<u>Average CO</u> <u>mole fraction</u>
Spring	1	1243	1853.4 ± 2.7	<u>185.4 ± 3.3</u>
	2	685	1852.6 ± 3.4	<u>147.8 ± 3.6</u>
	3	2231	1877.6 ± 2.5	<u>175.6 ± 3.0</u>
	4	1093	1850.5 ± 2.2	<u>196.0 ± 4.3</u>
	5	4108	1860.8 ± 1.5	<u>137.8 ± 1.1</u>
Summer	1	3981	1869.9 ± 1.2	<u>173.2 ± 2.4</u>
	2	2244	1878.5 ± 2.6	<u>135.0 ± 2.6</u>
	3	1040	1866.3 ± 2.6	<u>165.3 ± 5.2</u>
	4	916	1857.8 ± 2.4	<u>152.2 ± 5.1</u>
	5	578	1876.3 ± 5.1	<u>146.5 ± 5.4</u>
Autumn	1	1133	1870.1 ± 3.7	<u>159.1 ± 8.0</u>
	2	4235	1868.8 ± 1.3	<u>110.3 ± 1.3</u>
	3	2745	1873.6 ± 1.6	<u>135.1 ± 2.6</u>
	4	550	1865.2 ± 3.6	<u>149.1 ± 6.6</u>
Winter	1	3066	1879.8 ± 2.1	<u>159.9 ± 3.6</u>
	2	601	1872.6 ± 3.4	<u>282.2 ± 2.1</u>
	3	5261	1865.8 ± 1.0	<u>129.6 ± 1.7</u>

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1489 **Table 44.** Annual growth rates of atmospheric CH₄ in the City Regions (CR), the Tibetan Plateau
 1490 (TP), and ~~original~~ total regional records (~~Total~~) ~~during from~~ 1994 ~~to~~ -2019~~6~~ at ~~the~~ WLG station.

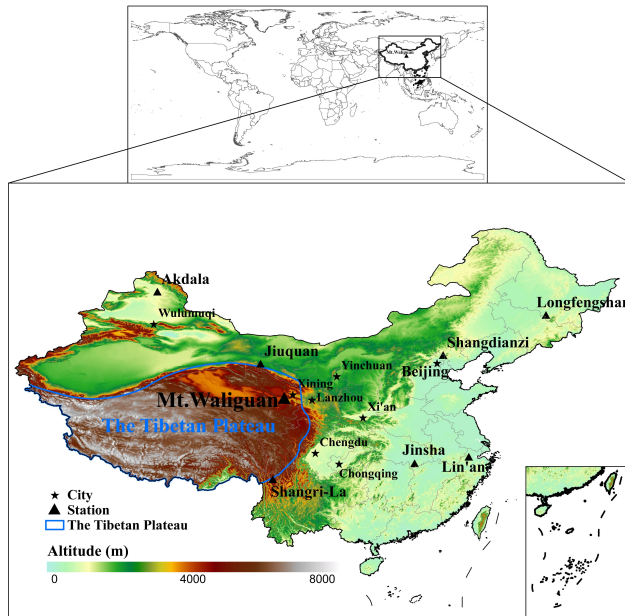
	1994-1997	1998-2002	2003-2007	2008-2012	2013- 2016 2019	1994- 2016 2019
CR	3.02 .8 ± 0.1	3.6 ± 0.2	5. 36 ± 0.2	7. 04 ± 0.2	6.25 .5 ± 0.1	5. 20 ± 0.1
TP	3.4 ± 0.1	3.0 ± 0.2	6. 78 ± 0.2	5. 76 ± 0.2	5.77 .3 ± 0.1	5. 12 ± 0.1
Total	4. 96 ± 0.1	2. 56 ± 0.2	4.95 .3 ± 0. 12	7. 76 ± 0. 12	5. 57 ± 0.1	5.1 ± 0.1

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1492 **Table 55.** The statistics of CH₄ data without air mass transport from city regions (TR) over different
 1493 periods during 2005-2016 at the WLG station.

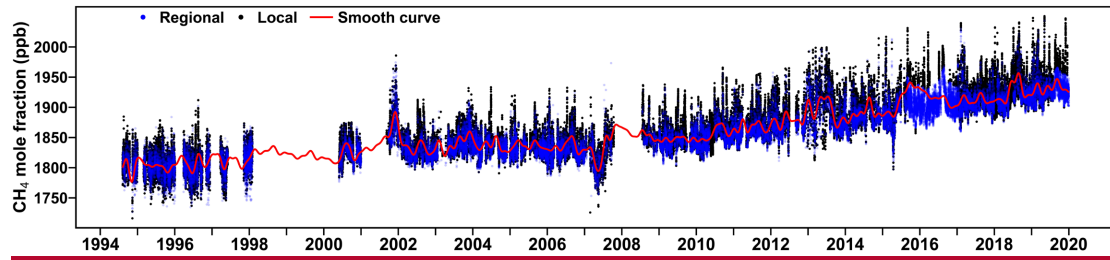
	Transport regions	Hours	Percentage (%)	Average (ppb)	Updated growth rate (ppb yr ⁻¹)
2005-2007	TR	6922	77.2	1824.9 ± 0.2	2.7 ± 0.2
	City	2041	22.8	1835.9 ± 0.5	-
2008-2012	TR	7060	64.4	1853.7 ± 0.2	10.1 ± 0.1
	City	4254	35.6	1861.0 ± 0.3	-
2013-2016	TR	4152	61.6	1888.2 ± 0.3	6.3 ± 0.1
	City	2591	38.4	1888.5 ± 0.5	-
2005-2016	TR	18134	67.1	1850.6 ± 0.2	7.0 ± 0.1
	City	8886	32.9	1863.2 ± 0.3	-

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1496 **Figure 1.** ~~The l~~ocation of Mt. Waliguan WMO/GAW global station as well as ~~the~~ other regional
 1497 stations in China. The gradient color indicates altitude. The digital elevation model (DEM) was
 1498 downloaded from Geospatial Data Cloud site, Computer Network Information Center, Chinese
 1499 Academy of Sciences (<http://www.gscloud.cn>), and then processed by ArcGis software. The China
 1500 map was derived from © National basic Geomatics Center of China (<http://www.ngcc.cn/ngcc/>).
 1501 The world map was obtained from © OpenStreetMap (<https://www.openstreetmap.org/>). ~~And the~~
 1502 ~~other shpfile file data and entire map were created by ArcGis software.~~



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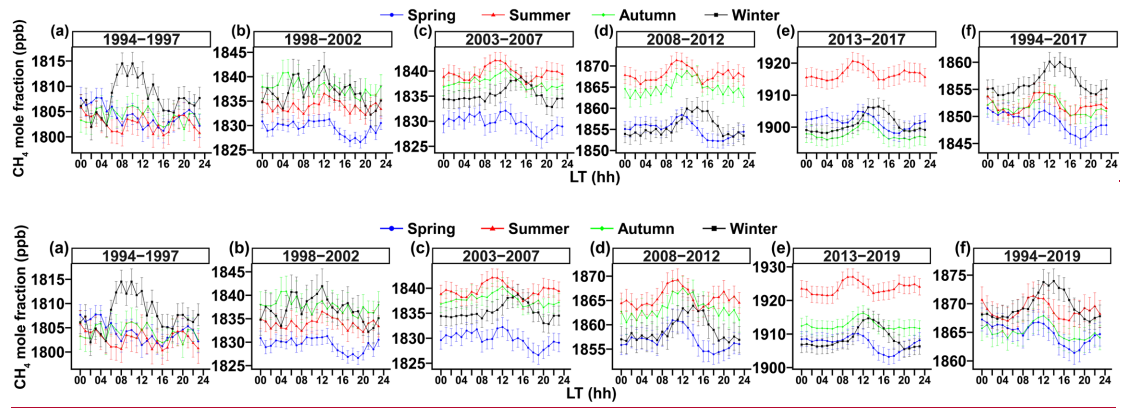
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Figure 2. The filtered hourly CH₄ data series from 1994 to 2019 at the WLG station. The transparent blue points are regionally representative data. The black points are locally influenced data. The red lines are smooth values of the regional data obtained by the curve-fitting routine of Thoning et al. (1989).

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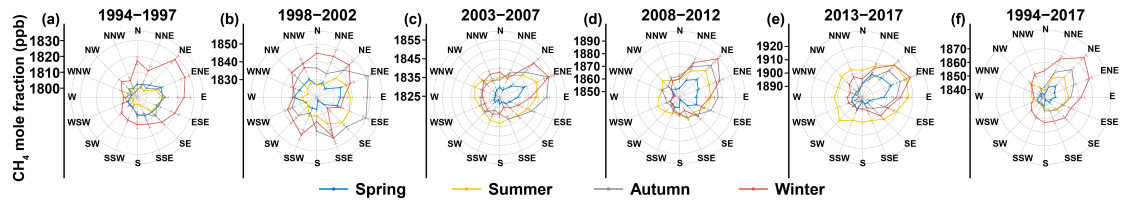
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1511 **Figure 3.** Diurnal CH₄ cycles in different periods from 1994 to ~~2017~~2019 at the WLG station. The

1512 lines with different colors represent various seasons. Error bars indicate the 95% confidence

1513 intervals.



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Figure 4. The wind-rose distribution of the average hourly CH_4 records mole fractions from 16

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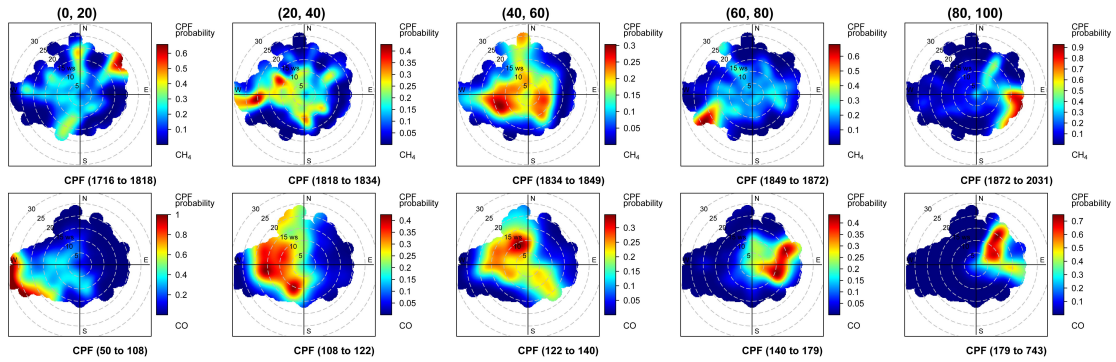
horizontal wind directions over different periods during 1994-2017 at the WLG station. The

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different colors represent the CH_4 data in different seasons. Error bars in all directions indicate 95%

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confidence intervals.



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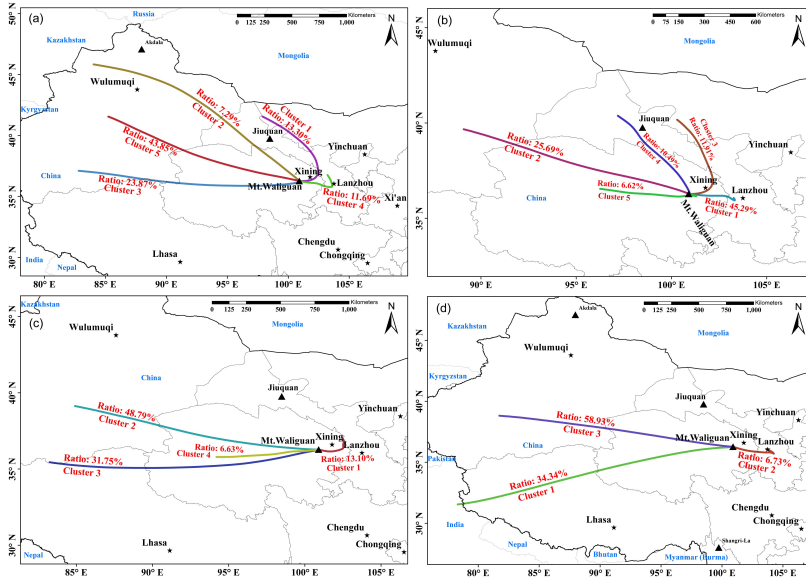
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Figure 5. The polar plot of the distribution of the CH₄ and CO concentration probability probabilities in different percentile ranges at the WLG station. The analysis was based on the conditional probability functions (CPF) by Ashbaugh et al. (1985). The top plots show the measurements-analysis of CH₄ from 1994 to 2017. The bottom plots show the CO measurements in 2004-2017. ‘ws’ means-refers to the wind speed. The values in-at the bottom of each panel show the range of concentrations in the relevant percentile range. Gradient colors represent the levels of CPF probability in different percentile ranges.



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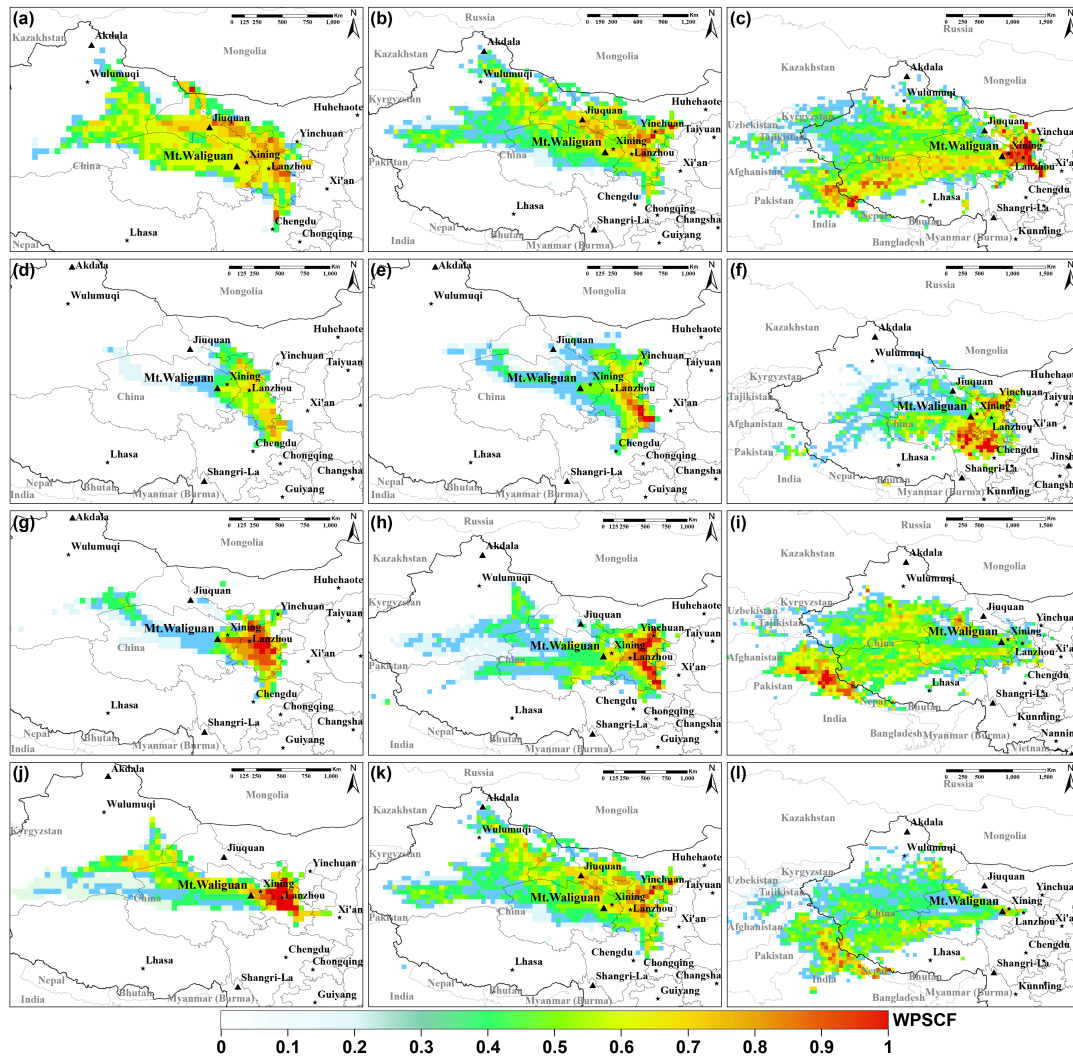
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Figure 6. Cluster analysis ~~to~~ of the 72-h back trajectories in different seasons (spring (a), summer (b), autumn (c), and winter (d)) during 2004-2017 ~~ending~~ at WLG station. ~~The (a), (b), (c) and (d) represents spring, summer, autumn and winter, respectively. The colored lines with different colors denoterepresent~~ different cluster ~~analysis results~~. The proportion of trajectories ~~in~~ each cluster is also marked.



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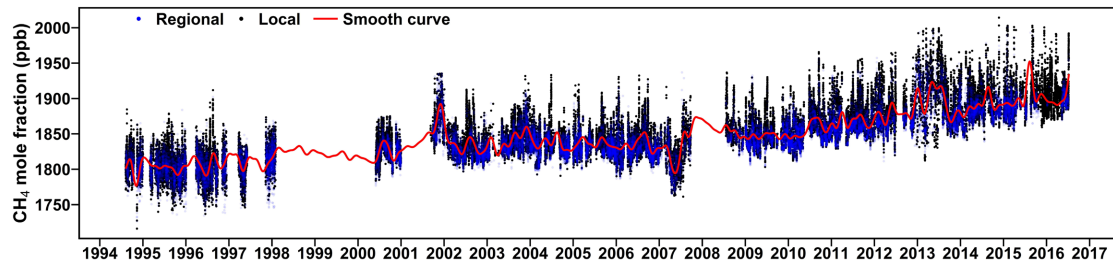
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Figure 7. The geographical distribution of the weighted potential sources of CH₄ in different periods over 1994-2017 at the WLG station. The gradient color shows the strong levels of potential source regions in different seasons, i.e. spring (a, b, c), summer (d, e, f), autumn (g, h, i), and winter (j, k, l), and different periods, i.e. 2004-2007 (a, d, g, j), 2008-2012 (b, e, h, k), and 2013-2017 (c, f, i, l).



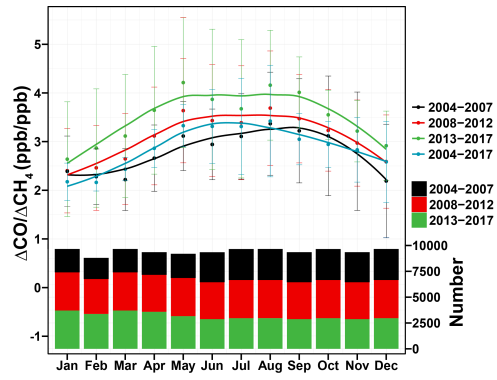
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Figure 7. Filtered hourly CH₄ data series from 1994 to 2017 at WLG station. The blue points with transparency represent regional events. The black points are the filtered local events. The red lines are calculated smooth values to the regional data by curve fitting routine of Thoning et al. (1989).



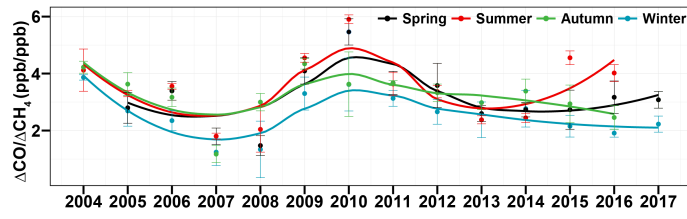
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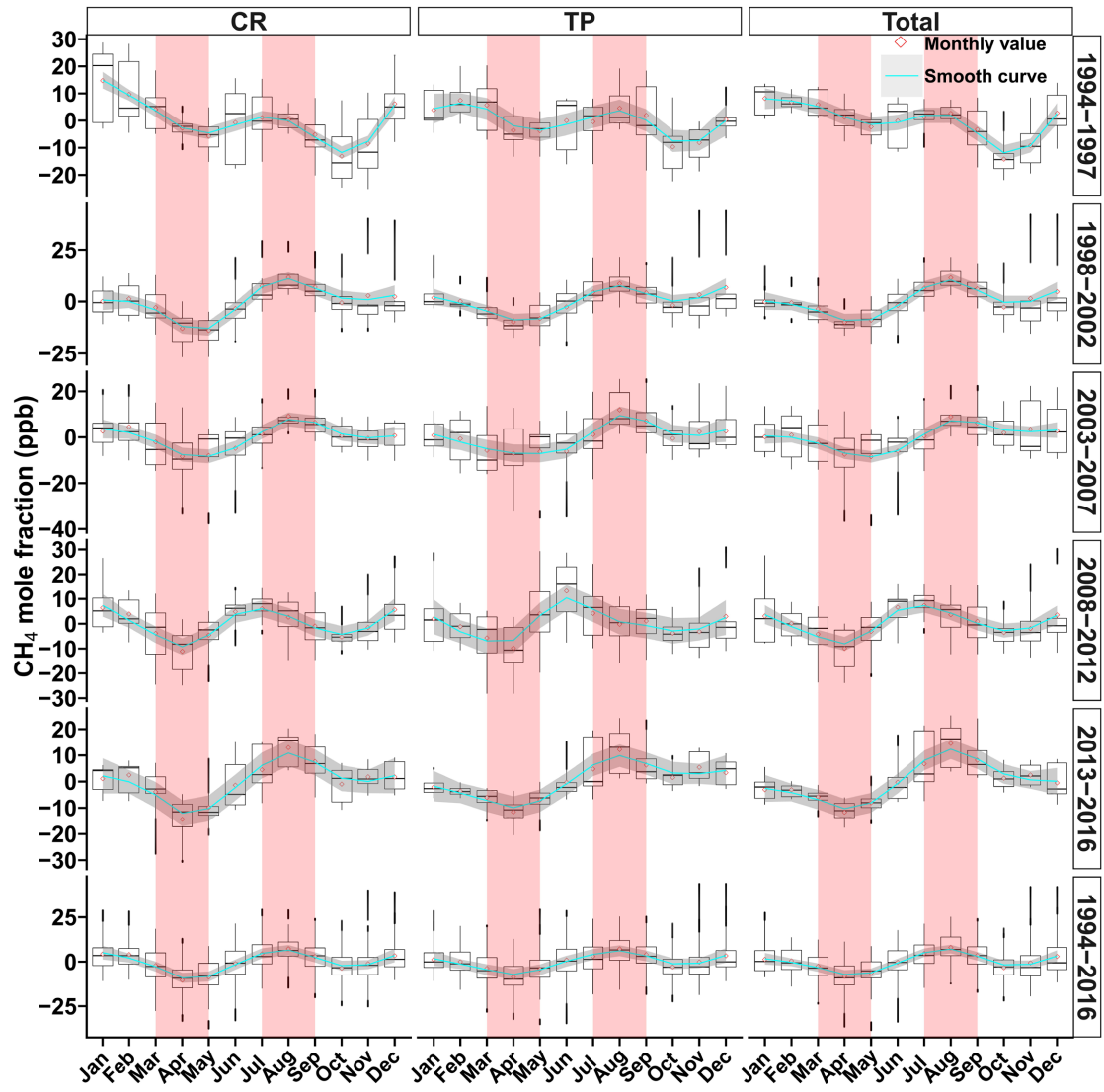
Figure 8. ~~A~~The average seasonal variation of the $\Delta\text{CO}/\Delta\text{CH}_4$ slopes in different periods ~~over~~ during 2004-2017 at the WLG station. The error bars show the standard deviation of the monthly averages. The vertical bars are the monthly numbers of data in different periods.

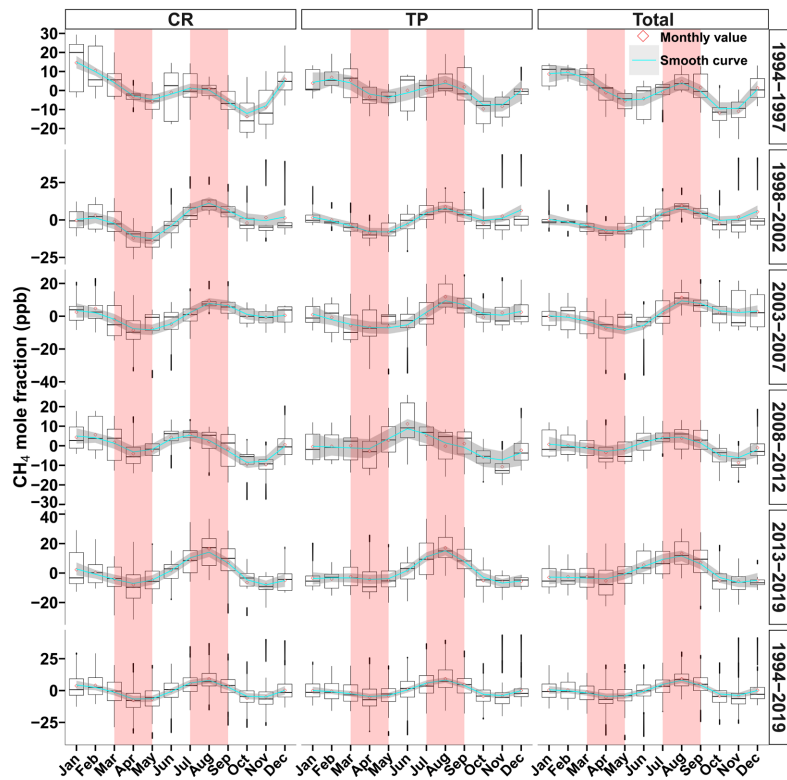


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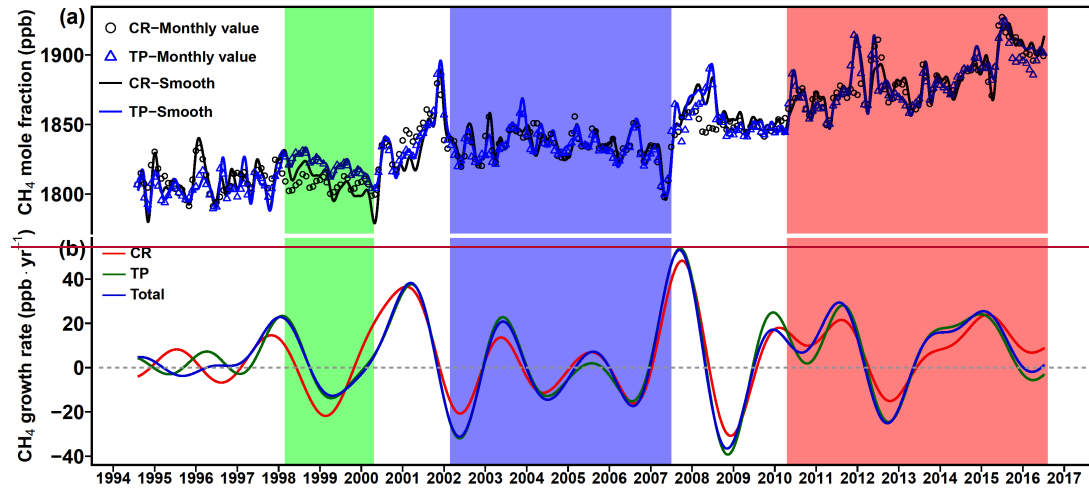
Figure 9. The long-term trend of $\Delta\text{CO}/\Delta\text{CH}_4$ slopes over 2004-2017 at WLG station.





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1551 **Figure 10.** Monthly variations of regional CH₄ mole fractions from 1994 to ~~2016~~2019 at the WLG
 1552 station. The ‘CR’, ‘TP’ and ‘Total’ represents the measurements from the City Regions, the Tibetan
 1553 Plateau and the ~~original~~-total regional records, respectively. The box ~~respectively~~ shows the 25th
 1554 percentile, the median and the 75th percentile from bottom to top. The bottom and the top whiskers
 1555 respectively reaches the minimum and 1.5 times the IQR (interquartile range). The black points are
 1556 identified as outliers. The red squares are the averages. The cyan lines are the smoothed curve of
 1557 the averages using the method of loess (Local Polynomial Regression Fitting). The gray bands are
 1558 the 95% confidence interval of smoothed curve.



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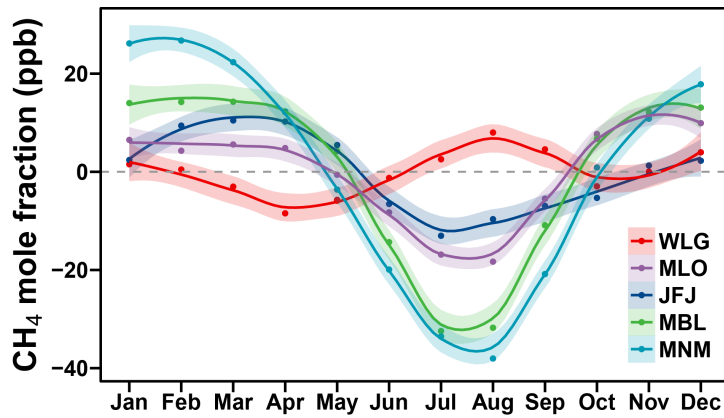
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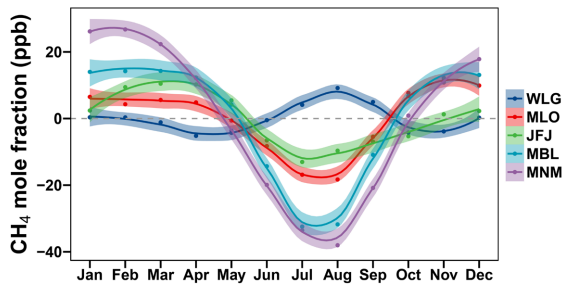
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Figure 11. The top panel (a) shows the smooth curve and monthly values of CH₄ mole fraction in the City Regions (CR) and the Tibet Plateau (TP) during 1994–2016 at WLG station. The bottom panel (b) is the annual growth rates of atmospheric CH₄ records from CR, TP as well as the total regional time series (Total). The growth rates are calculated from the first derivative of trend curves. The smooth curve and the trend is calculated by the method of Thoning et al. (1989).

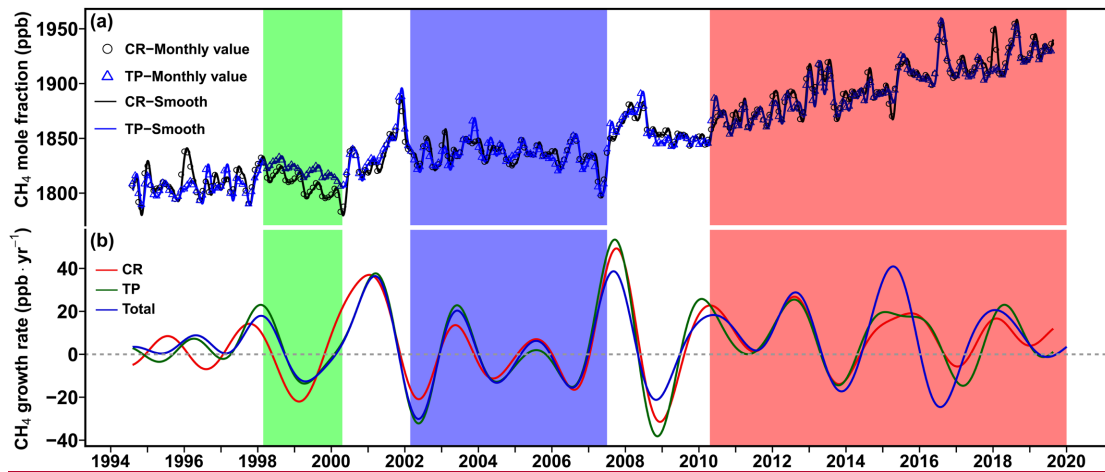


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1567 **Figure 11-12.** The seasonal cycles of atmospheric CH₄ observed ~~over 1994-2017~~ at the WMO/GAW
 1568 global stations of Mauna Loa (MLO, 1994-2018), Jungfraujoch (JFJ, 2005-2018), Minamitorishima
 1569 (MNM, 1994-2019), and Mt. Waliguan (WLG, 1994-2019) in the Northern Hemisphere. The data
 1570 of other sites (except WLG) were from WDCGG. The data in the marine boundary layer (MBL,
 1571 1994-2019) were from NOAA / ESRL lab at the a similar latitude to the WLG. _____



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Figure 12. The top panel (a) shows the smoothed curves and monthly values of the CH₄ mole fractions in the City Regions (CR) and the Tibetan Plateau (TP) during 1994-2019 at the WLG station. The bottom panel (b) is the annual growth rates of the atmospheric CH₄ records from the CR and the TP as well as the total regional time series. The growth rates were calculated from the first derivative of the trend curves. The smoothed curves and the trends were calculated using the method of Thoning et al. (1989).