

## Comments to the Author:

The authors have reasonably well addressed the comments of the two anonymous referees and they have modified their manuscript accordingly. However, several alterations and corrections are needed for the Main text and Supplement before the manuscript can be published in ACP.

**Response:** Thank you for providing us the opportunity to revise and improve our manuscript. The comments on the main text and supplement are addressed accordingly. Detailed responses to the comments are provided in blue. Attached please also find the marked-up manuscript to track the changes in the revised manuscript.

### 1. Comments on the main text and supplement

#### Main text:

Line 496: Replace "201613401" by "114, E1054-E1061, <https://doi.org/10.1073/pnas.16130114>".

**Response:** We add the article number, page ranges and doi number:

“Winiger, P., Andersson, A., Eckhardt, S., Stohl, A., Semiletov, I. P., Dudarev, O. V., Charkin, A., Shakhova, N., Klimont, Z., and Heyes, C.: Siberian Arctic black carbon sources constrained by model and observation, P. Natl. Acad. Sci. USA, 114, E1054-E1061, <https://doi.org/10.1073/pnas.1613401114>, 2017.”

Lines 555-557: Several of the abbreviations in the abscissa of the Figure are only defined in Table S1. Therefore, the Figure caption should be extended and reference should be made in it to Table S1.

**Response:** Thank you for this point. The revised Figure caption shows (page 24):

“**Figure 6.** Fraction of fossil carbon in EC, OC and SOC ( $f_{\text{fossil}}(\text{EC})$ ,  $f_{\text{fossil}}(\text{OC})$  and  $f_{\text{fossil}}(\text{SOC})$ , respectively) during haze and clean periods in Xi'an and Beijing. Interquartile ranges (25th–75th percentile) of the median  $f_{\text{fossil}}(\text{SOC})$  are shown as vertical bars in purple. Uncertainties of  $^{14}\text{C}$ -apportioned  $f_{\text{fossil}}(\text{EC})$  and  $f_{\text{fossil}}(\text{OC})$  are indicated but are too small to be visible. The studied samples include day and night samples in Xi'an (X) and 24 h integrated samples in Beijing (B) during haze periods (“H” samples) and clean periods (“C” samples). For details of each sample (x-axis), see Fig. 1 and Table S1.”

We also make reference to Fig. 1, because Fig.1 has the same x-axis (i.e., sample name) as Fig.6. In Fig.1, the definition of each sample name is indicated in the plot area of the Figure.

#### Supplement:

Page S2, line 5 from bottom: Abbreviations and acronyms, here "PDFs", should be defined (written full-out) when first used. "PDFs" is only defined on page S3.

**Response:** In the revised Supplement, we present both the full name and abbreviation of “PDFs” when first used (page S2). After the abbreviation is defined, only “PDFs” is used (e.g., page S3).

**2. Response:** Alterations and corrections are made following the editor's comments. Below are the alterations made in the main text and supplement.

#### For the Main text:

Line 2: Replace "reveal" by "reveals".

Line 49: Replace "quantification" by "quantifying".

Line 52: Replace "provide" by "provides".

Line 93: Replace "IMPROVE\_A" by "The IMPROVE\_A".

Line 94: Replace "measurements" by "measurement".

Line 116: Replace "is smaller" by "are smaller".

Line 158: Replace "equal to" by "are equal to".

Line 163: Abbreviations and acronyms, here "PDFs", should be defined (written full-out) when first used. Since "PDFs" seems to stand for "probability density functions" and is not used elsewhere in the Main text, it should simply be replaced here by "probability density functions".

Line 177: Replace "is resulted from" by "results from the".

Line 210: Replace "vary spatially" by "varies spatially".

Line 230: Replace "sources," by "sources;".

Line 254: Replace "e.g.," by "i.e.,".

Line 265: Replace ", in contrast to" by "; in contrast".

Line 269: Replace ", mainly" by "; this mainly".

Line 318: Abbreviations and acronyms, here "SOA", should be defined (written full-out) when first used. Since "SOA" stands for "secondary organic aerosol" and is not used elsewhere in the Main text, it should simply be replaced here by "secondary organic aerosol".

Line 327: Replace "governs the" by "govern the".

Line 339: Replace "dominate fossil" by "dominating fossil".

### **For the Supplement:**

Page S1, line 3: Replace "reveal" by "reveals".

Page S2, line 2: Replace "associate" by "associated".

Page S2, line 6: Replace "agrees each" by "agree with each".

Page S15, line 3: The abbreviated journal name should be in regular font instead of in italic.

### **3. Besides adjustments made above, we checked the main text and Supplement, and found that several changes are needed:**

#### **Main text:**

**(a). Page 3, line 90 (the revised version).** The end of sampling date in Beijing should be 8 January 2017 (it was a written mistake as 7 January 2017 in the original manuscript), as shown in Fig. S1 and Table S1.

“In Beijing, the 24 h integrated (10:00 a.m. to 10:00 a.m. the following day) PM<sub>2.5</sub> was collected from 2 December 2016 to 8 January 2017.”

**(b). Page 7, line 208.** The standard deviation (SD) of  $f_{\text{fossil}}(\text{OC})$  in Xi'an should be 2%, not 3% in the original manuscript. The SD can be calculated from  $f_{\text{fossil}}(\text{OC})$  values in Xi'an in Table S3. This correction will not affect any conclusion from this study. Sorry that we were not aware this typo before.

“The presented overall average  $f_{\text{fossil}}(\text{OC})$  for winter 2016/2017 in Beijing ( $66 \pm 3\%$ ) was higher than that in Xi’an ( $46 \pm \underline{2}\%$ )”

**Supplement:**

**(c). Title of Table S4 (page S12).** We add the abbreviation ( $\text{POC}_{\text{fossil}}$ ) for primary OC from fossil sources in the title:

“**Table S4.** Concentrations ( $\mu\text{g m}^{-3}$ ) of primary OC from biomass burning ( $\text{POC}_{\text{bb}}$ ), primary OC from fossil sources ( $\text{POC}_{\text{fossil}}$ )...”

**(d). Title of Table S5 (page S13).** It should be “in different cities”, not “in different seasons”

“**Table S5.** Fractional contribution of different incomplete combustion sources to EC in different cities...”

# Measurement report: Dual-carbon isotopic characterization of carbonaceous aerosol reveals different primary and secondary sources in Beijing and Xi'an during severe haze events

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**Abstract.** To mitigate haze pollution in China, a better understanding of the sources of carbonaceous aerosols is required due to the complexity in multiple emissions and atmospheric processes. Here we combined the analysis of radiocarbon and the stable isotope <sup>13</sup>C to investigate the sources and formation of carbonaceous aerosols collected in two Chinese megacities (Beijing and Xi'an) during severe haze events of “red alarm” level from December 2016 to January 2017. The haze periods with daily PM<sub>2.5</sub> concentrations as high as ~400 µg m<sup>-3</sup> were compared to subsequent clean periods (i.e., PM<sub>2.5</sub> < median concentrations during the winter 2016/2017), with PM<sub>2.5</sub> concentrations below 100 µg m<sup>-3</sup> in Xi'an and below 20 µg m<sup>-3</sup> in Beijing. In Xi'an, liquid fossil fuel combustion was the dominant source of elemental carbon (EC; 44%–57%), followed by biomass burning (25%–29%) and coal combustion (17%–29%). In Beijing, coal combustion contributed 45%–61% of EC and biomass burning (17%–24%) and liquid fossil fuel combustion (22%–33%) contributed less. Non-fossil sources contributed 51%–56% of organic carbon (OC) in Xi'an and fossil sources contributed 63%–69% of OC in Beijing. Secondary OC (SOC) was largely contributed by non-fossil sources in Xi'an (56 ± 6%) and by fossil sources in Beijing (75 ± 10%), especially during haze periods. The fossil vs. non-fossil contributions to OC and EC did not change drastically during haze events in both Xi'an and Beijing. However, compared to clean periods, the contribution of coal combustion to EC during haze periods increased in Xi'an and decreased in Beijing. During clean periods, primary OC from biomass burning and fossil sources constituted ~70% of OC in Xi'an and ~53% of OC in Beijing. From clean to haze periods, the contribution of SOC to total OC increased in Xi'an, but decreased in Beijing, suggesting that contribution of secondary organic aerosol formation to increased OC during haze periods was more efficient in Xi'an than in Beijing. In Beijing, the high SOC fraction in total OC during clean periods was mainly due to elevated contribution from non-fossil SOC. In Xi'an, a slight day-night difference was observed during the clean period, with enhanced fossil contributions to OC and EC during the day. This day-night difference was negligible during severe haze periods, likely due to enhanced accumulation of pollutants under stagnant weather conditions.

## 1 Introduction

Severe haze pollution with high PM<sub>2.5</sub> (i.e., particulate matter with aerodynamic diameter  $\leq 2.5 \mu\text{m}$ ) concentrations and reduced visibility occurs frequently during winter in China (An et al., 2019). Field measurements show that carbonaceous aerosol contributes a significant fraction of PM<sub>2.5</sub> loading during severe haze events in China (Huang et al., 2014; Elser et al., 2016; Liu et al., 2016). Therefore, a better understanding of the sources and atmospheric processes of carbonaceous aerosols is needed for mitigating haze pollution. Many previous studies focus solely on Beijing, the capital of China. However, studies on other megacities are also needed for comparison as well as for a more comprehensive understanding of haze pollution in China.

Carbonaceous aerosol constituents are separated into elemental carbon (EC) and organic carbon (OC), fractions differing in their thermal refractiveness with EC being thermally refractory and OC weakly refractory (Pöschl, 2003, 2005; Petzold et al., 2013). EC is emitted as primary particles from incomplete combustion sources (i.e., biomass burning and fossil fuel combustion). Unlike EC, OC can either be emitted as primary OC (POC) from combustion sources and non-combustion sources (e.g., biogenic emissions) or formed in the atmosphere as secondary OC (SOC) via the reaction of gas precursors (Hallquist et al., 2009; Jimenez et al., 2009). The sources and abundance of different carbon fractions in carbonaceous aerosols vary considerably in different Chinese cities, as a result of complex interplay between meteorology, local and regional emissions sources, and atmospheric processes (Zhang et al., 2008; Cui et al., 2015; Tie et al., 2017; An et al., 2019). Therefore, [quantification-quantifying](#) the sources of carbonaceous aerosol in China is a challenging task.

Radiocarbon ( $^{14}\text{C}$ ) analysis of carbonaceous aerosols is the most direct and effective method to distinguish their main sources, exploiting the fact that OC and EC of fossil origins (i.e., vehicle emissions, coal combustion) do not contain  $^{14}\text{C}$  (Heal, 2014; Cao et al., 2017; Dusek et al., 2013).  $^{14}\text{C}$  analysis of OC and EC separately provides a clear-cut division of carbonaceous aerosols into four major fractions: fossil OC, non-fossil OC (e.g., OC from biomass burning, biogenic emissions and cooking), fossil EC and biomass-burning EC (e.g., Gustafsson et al., 2009; Szidat et al., 2009; Zotter et al., 2014; Dusek et al., 2017; Ni et al., 2018, 2019a). For example, Liu et al. (2014) demonstrated that fossil sources including coal burning and vehicle emissions dominated EC during winter haze events in Guangzhou, China. Zhang et al. (2015) showed that the elevated carbonaceous aerosols during the severe haze event in January 2013 in China were by a large extent driven by SOC from both fossil and non-fossil precursors. In addition, the analysis of the  $^{13}\text{C}/^{12}\text{C}$  ratio can refine  $^{14}\text{C}$  source apportionment, because coal combustion and vehicle emissions have different  $^{13}\text{C}$  source signatures although they both completely depleted in  $^{14}\text{C}$  (e.g., Andersson et al., 2015; Li et al., 2016; Winiger et al., 2016, 2017; Fang et al., 2017, 2018; Ni et al., 2018).

A critical question for effective haze mitigation is whether carbonaceous aerosols in different Chinese cities have similar characteristics during haze events. However, there are not many studies highlighting the differences in sources of primary and secondary carbonaceous aerosols between cities, especially for studies employing the analysis of  $^{14}\text{C}$  or the stable isotope  $^{13}\text{C}$  (e.g., Zhang et al., 2015; Andersson et al., 2015; Liu et al., 2016). In this work, we compare the severe haze events reaching

65 “red alarm” level (i.e., the highest air-quality warning level in China) in two Chinese megacities (Beijing and Xi’an) during  
December 2016 and January 2017. We present measurements of dual carbon isotopes (i.e.,  $^{14}\text{C}$  and the stable carbon isotope  
 $^{13}\text{C}$ ) for EC and OC. The sources of carbonaceous aerosols are elucidated and compared between haze and clean periods in  
Beijing and Xi’an, with the main objectives: (1) quantitative understanding of the difference in EC contribution from burning  
of biomass, coal and liquid fossil fuel (i.e., vehicle emissions) under different pollution conditions; and (2) constraint on the  
70 sources of both primary and secondary OC. Furthermore, the comparison of day-time and night-time results in Xi’an yields  
insight into diurnal variation in sources of carbonaceous aerosols.

## 2 Methods

### 2.1 Aerosol collection

To collect  $\text{PM}_{2.5}$  samples, high-volume aerosol samplers (flow rate =  $1.0 \text{ m}^3 \text{ min}^{-1}$ ; TE-6070 MFC, Tisch Inc., Cleveland, OH,  
75 USA) were used at an urban background site in Xi’an and Beijing (see Table S1 for details about the sampling sites). Xi’an is  
the largest city in northwestern China, with over 8.8 million residents and 2.5 million vehicles in 2016 (Xi’an Municipal Bureau  
of Statistics and NBS Survey Office in Xi’an, 2017). Surrounded by Qinling Mountains to the south and the Loess Plateau to  
the north, days with low wind speed occur frequently in Xi’an, promoting the accumulation of air pollutants. Xi’an is now  
facing increased serious air quality issues due to the rapid increase of motor vehicles and energy consumption in the past two  
80 decades. Besides residential coal combustion, biomass burning is also a major emission source in Xi’an and its surrounding  
areas (i.e., Guanzhong Plain) for heating and cooking especially in winter (Zhang et al., 2014; Xu et al., 2016). Beijing, the  
capital of China, is a megacity with over 21 million residents and 5.7 million vehicles in 2016 (Beijing Municipal Bureau of  
Statistics and NBS Survey Office in Beijing, 2017). Beijing is located in the Beijing-Tianjin-Hebei region, the most  
economically developed region in North China. However, the rapid economic growth and urbanization associated with heavy  
85 coal consumption and rapid increase usage of vehicles lead to the poor air quality in Beijing. Besides local emissions, regional  
transport of pollutants between neighboring cities also contributes to air pollution in Beijing (Zheng et al., 2015; An et al.,  
2019). The 12 h integrated (daytime: 8:00 a.m. to 8:00 p.m., local standard time, LST; nighttime: 8:00 p.m. to 8:00 a.m. the  
following day)  $\text{PM}_{2.5}$  was sampled on pre-combusted quartz filters (8 in  $\times$  10 in; QM-A, Whatman Inc., Clifton, NJ, USA) in  
Xi’an from 1 January 2017 to 10 January 2017. In Beijing, the 24 h integrated (10:00 a.m. to 10:00 a.m. the following day)  
90  $\text{PM}_{2.5}$  was collected from 2 December 2016 to 7–8 January 2017. Field blanks were collected by exposing filters to ambient  
air for 15 min. Immediately after collection, the filters were transferred into pre-combusted aluminum foils and stored  
at  $-18^\circ\text{C}$ .

### 2.2 Concentration measurements of OC and EC

The IMPROVE\_A protocol (Chow et al., 2007) was implemented on a carbon analyzer (DRI Model 2001, Atmoslytic Inc.,  
95 USA) for measurements of carbon concentrations. The relative standard deviations for the replicate analyses were smaller than

10 % for OC and EC. OC mass was corrected for field blanks ( $0.4 \mu\text{g cm}^{-2}$ ). EC was too small to be detected on field blanks. Acidification to remove potential interferences from carbonates is not necessary, because carbonate carbon in  $\text{PM}_{2.5}$  samples is found to be negligible, compared to the relatively larger OC and EC amounts for both mass determination and carbon isotopic analysis (Supplement S1).

## 100 2.3 Analysis of carbon isotope

Six samples from haze and clean days were selected per sampling site for carbon isotope analysis (Tables S1 and S2). We define clean days at each site as  $\text{PM}_{2.5} < \text{median concentration in the winter heating season from 15 November 2016 to 15 March 2017}$ . In Xi'an, there were 4 composite samples (2 daytime + 2 nighttime) from haze days, and 2 composite samples (1 daytime + 1 nighttime) from clean days. In Beijing, five 24 h samples were selected from haze days, and 1 composite sample  
105 from two clean days. Each composite sample consists of 2 12h (for Xi'an) or 24 h (for Beijing) filter pieces with similar  $\text{PM}_{2.5}$  loadings that agree within 20 % (Fig. S1).

### 2.3.1 Stable isotope $^{13}\text{C}$

Filter samples were placed in a quartz tube with CuO grains. The tube was subsequently evacuated and sealed before heating for 3 h at  $375^\circ\text{C}$  to remove OC. Then the EC was extracted by heating the remaining carbon for 5 h at  $850^\circ\text{C}$ . The  $^{13}\text{C}/^{12}\text{C}$   
110 ratio of EC was measured by an isotope mass spectrometer (Finnigan MAT-251; Bremen, Germany) and expressed in the delta notation:

$$\delta^{13}\text{C} = \left[ \frac{(^{13}\text{C}/^{12}\text{C})_{\text{sample}}}{(^{13}\text{C}/^{12}\text{C})_{\text{V-PDB}}} - 1 \right]. \quad (1)$$

$\delta^{13}\text{C}$  values are usually reported in per mil (‰).  $(^{13}\text{C}/^{12}\text{C})_{\text{V-PDB}}$  is the  $^{13}\text{C}/^{12}\text{C}$  ratio of the international standard Vienna Pee Dee Belemnite (V-PDB). A well-characterized standard was measured every working day. Duplicate analysis of  $\delta^{13}\text{C}$  of EC showed  
115 an analytical precision better than  $\pm 0.3\text{‰}$ . This method was detailed in Ni et al., (2019b), where impacts of potential charred OC on the isolated EC were evaluated using an isotope-mass-balance based sensitivity analysis. We concluded that the expected differences in  $\delta^{13}\text{C}_{\text{EC}}$  ~~is-are~~ smaller than 1‰ under the assumption that the fraction of charred OC in the isolated EC is at most 20%.

### 2.3.2 Radiocarbon

120 OC and EC in  $\text{PM}_{2.5}$  samples were converted to  $\text{CO}_2$  using an aerosol combustion system (ACS; Dusek et al., 2014). The ACS has been evaluated in two intercomparison studies (Szidat et al., 2013; Zenker et al., 2017). The isolated  $\text{CO}_2$  was subsequently reduced to graphite (de Rooij et al., 2010) before  $^{14}\text{C}$  measurements can be conducted with the accelerator mass spectrometer (AMS) at CIO (van der Plicht et al., 2000). The temperature protocol for OC and EC combustion has been detailed in Ni et al. (2018), and is summarized in Fig. S2. To remove possible interfering gas (e.g.,  $\text{NO}_x$ , halogen and water vapor) from  $\text{CO}_2$ , a

125 reduction oven filled with copper grains and silver, a dry ice-ethanol bath and a flask filled with phosphorus pentoxide are installed on the ACS.

Fraction modern ( $F^{14}\text{C}$ ) is used to report the  $^{14}\text{C}$  data (Reimer et al., 2004).  $F^{14}\text{C}$  relates the  $^{14}\text{C}/^{12}\text{C}$  ratio of a sample to the ratio of the unperturbed atmosphere in the reference year 1950:

$$F^{14}\text{C} = \frac{(^{14}\text{C}/^{12}\text{C})_{\text{sample}, [-25]}}{(^{14}\text{C}/^{12}\text{C})_{1950, [-25]}}. \quad (2)$$

130 Both ratios are normalized to  $\delta^{13}\text{C}$  of -25‰ to remove the effect of isotope fractionation. Practically,  $(^{14}\text{C}/^{12}\text{C})_{1950, [-25]}$  equals to the  $^{14}\text{C}/^{12}\text{C}$  ratio of an oxalic acid standard (OXII) multiplied by a factor of 0.7459. Contamination during graphitization and AMS measurements was quantified from the measured  $F^{14}\text{C}$  of standards (OXII with known  $F^{14}\text{C}$  of 1.3407 and Rommenhöller with  $F^{14}\text{C}=0$ ) processed in the same way as samples. The resulting estimated dead and modern contamination were used to correct the  $^{14}\text{C}$  data according to Santos et al. (2007). The reliability of data correction was further verified by  
135 measuring two secondary standards (i.e., IAEA-C7 and-C8) on the same wheel of samples. The measured values of IAEA-C7 ( $0.495 \pm 0.008$ ) and IAEA-C8 ( $0.154 \pm 0.007$ ) agree with their respective consensus value ( $0.4953 \pm 0.0012$  and  $0.1503 \pm 0.0017$ ) within uncertainties.

## 2.4 Source apportionment methods

$F^{14}\text{C}$  is larger than the fraction of non-fossil carbon (i.e.,  $f_{\text{nf}}(\text{OC})$  for OC,  $f_{\text{bb}}(\text{EC})$  for EC) due to the large release of  $^{14}\text{C}$  into  
140 the atmosphere from the nuclear bomb tests in 1960s. To eliminate this effect,  $F^{14}\text{C}$  is divided by  $F^{14}\text{C}$  of non-fossil sources ( $F^{14}\text{C}_{\text{nf}}$ ).  $F^{14}\text{C}_{\text{nf}}$  is estimated as  $1.09 \pm 0.05$  for OC and  $1.10 \pm 0.05$  for EC (see details in Ni et al., 2019b), using a tree growth model and the contemporary atmospheric  $^{14}\text{CO}_2$  over the past years (Lewis et al., 2004; Mohn et al., 2008; Levin et al., 2010), with the assumption that biomass-burning OC and biogenic OC contribute to 85% and 15% of total OC, respectively. Once  $f_{\text{nf}}(\text{OC})$  and  $f_{\text{nf}}(\text{EC})$  are known, carbon concentrations can be apportioned into EC and OC from non-fossil sources ( $\text{EC}_{\text{bb}}$ ,  $\text{OC}_{\text{nf}}$ )  
145 and fossil sources ( $\text{EC}_{\text{fossil}}$ ,  $\text{OC}_{\text{fossil}}$ ) (Eq. 3–6 in Table 1).  $\text{OC}_{\text{nf}}$  and  $\text{OC}_{\text{fossil}}$  are further divided into POC from biomass burning ( $\text{POC}_{\text{bb}}$ ), other non-fossil OC ( $\text{OC}_{\text{o,nf}}$ ) (Eq. 7–8), primary and secondary fossil OC ( $\text{POC}_{\text{fossil}}$  and  $\text{SOC}_{\text{fossil}}$ , respectively; Eq. 9–10).  $\text{POC}_{\text{bb}}$  and  $\text{POC}_{\text{fossil}}$  are estimated using EC as a tracer of primary emissions (i.e., the EC tracer method; Turpin and Huntzicker, 1995). Based on  $\text{OC}_{\text{o,nf}}$  and  $\text{SOC}_{\text{fossil}}$ , total SOC and the fraction of fossil carbon in SOC ( $f_{\text{fossil}}(\text{SOC})$ ) are estimated using Eq. (11–12).  $\text{OC}_{\text{o,nf}}$  mainly includes SOC of non-fossil origins ( $\text{SOC}_{\text{nf}}$ ), primary biogenic OC and cooking OC.  $\text{OC}_{\text{o,nf}}$  is  
150 approximately  $\text{SOC}_{\text{nf}}$ , as contributions of primary biogenic sources and cooking to  $\text{OC}_{\text{o,nf}}$  are likely small (Hu et al., 2010; Guo et al., 2012). If cooking is prominent,  $\text{OC}_{\text{o,nf}}$  is an overestimate of  $\text{SOC}_{\text{nf}}$ . To estimate the uncertainties of the source apportionment results, a Monte Carlo simulation ( $n=10000$ ) using Eq. (3–12) was carried out as described in Supplement S3. The  $^{14}\text{C}$  source apportionment results are presented in Tables S3 and S4.



The dual carbon isotope signatures of EC were used in a Bayesian Markov chain Monte Carlo (MCMC) scheme (Andersson, 2011), to conduct the mass-balance three source apportionment of EC (e.g., Andersson et al., 2015; Li et al., 2016; Winiger et al., 2016, 2017; Fang et al., 2017, 2018). That is, the  $F^{14}C$  and  $\delta^{13}C$  of ambient EC ( $F^{14}C_{(EC)}$  and  $\delta^{13}C_{EC}$ ) can be explained by burning of biomass (bb), coal (coal) and liquid fossil fuel (liq.fossil; i.e., vehicle emissions):

$$\begin{pmatrix} F^{14}C_{(EC)} \\ \delta^{13}C_{EC} \\ 1 \end{pmatrix} = \begin{pmatrix} F^{14}C_{nf} & F^{14}C_{coal} & F^{14}C_{liq.fossil} \\ \delta^{13}C_{bb} & \delta^{13}C_{coal} & \delta^{13}C_{liq.fossil} \\ 1 & 1 & 1 \end{pmatrix} \begin{pmatrix} f_{bb} \\ f_{coal} \\ f_{liq.fossil} \end{pmatrix} \quad (13)$$

$F^{14}C_{coal}$  and  $F^{14}C_{liq.fossil}$  are equal to zero since coal and liquid fossil fuel do not contain  $^{14}C$ .  $\delta^{13}C_{bb}$ ,  $\delta^{13}C_{coal}$  and  $\delta^{13}C_{liq.fossil}$  are  $\delta^{13}C$  signatures for EC from the three sources. Their values were established as  $\delta^{13}C_{bb}$  ( $-26.7 \pm 1.8$  ‰ for C3 plants, and  $-16.4 \pm 1.4$  ‰ for corn stalk; mean  $\pm$  SD),  $\delta^{13}C_{coal}$  ( $-23.4 \pm 1.3$  ‰) and  $\delta^{13}C_{liq.fossil}$  ( $-25.5 \pm 1.3$  ‰), based on critical evaluations of literature studies (Andersson et al., 2015; Ni et al., 2018; and references therein). Uncertainties in  $F^{14}C$  and  $\delta^{13}C$  source signatures and the measured  $F^{14}C_{(EC)}$  and  $\delta^{13}C_{EC}$  are considered in the MCMC technique (Parnell et al., 2010, 2013). MCMC outputs are the posterior [PDFs probability density functions](#) for  $f_{bb}$ ,  $f_{coal}$  and  $f_{liq.fossil}$  (i.e., the relative contribution of each source to EC). The median and interquartile range (25th–75th percentile) are used as the best estimate and the uncertainties, respectively.

### 3 Results and discussion

#### 3.1 Fossil and non-fossil contributions to EC and OC

During the measurement periods, the highest daily mass concentrations of  $PM_{2.5}$  in Xi'an ( $\sim 250$ – $420 \mu g m^{-3}$ ) and Beijing ( $\sim 210$ – $360 \mu g m^{-3}$ ; Fig. S1) were 10–17 and 8–14 times higher than the standard of World Health Organization ( $25 \mu g m^{-3}$ ; WHO, 2006), respectively. Using radiocarbon measurements, we investigated the sources of carbonaceous aerosols in  $PM_{2.5}$  in both cities during several haze periods, and compared them to clean periods, with  $PM_{2.5}$  concentrations below  $100 \mu g m^{-3}$  in Xi'an and below  $20 \mu g m^{-3}$  in Beijing. In Xi'an, even during clean periods we defined here, daily  $PM_{2.5}$  concentrations were higher than the Chinese pollution standard of  $75 \mu g m^{-3}$ , reflecting severe air quality problems.  $PM_{2.5}$ , OC and EC concentrations during haze periods were  $> 2$  times higher in Xi'an and  $> 5$  times higher in Beijing than those during clean periods, respectively. OC/EC ratios in Xi'an slightly decreased from  $\sim 4$  during haze periods to  $\sim 3$  during clean periods, while OC/EC ratios in Beijing were lower during haze periods ( $\sim 3$ ) than clean periods ( $\sim 4$ ). This reflects different sources and formation mechanisms of haze pollution in the two cities. In Xi'an, we collected day and night  $PM_{2.5}$  samples. No consistent day-night variations in concentrations of  $PM_{2.5}$ , OC and EC (Figs. 1 and S1) were observed. This ~~is resulted results~~ from [the](#) diurnal cycle of human activities (e.g., traffic, usage of biomass and coal for heating or cooking) and the development of planetary boundary layer height which controls the vertical mixing and dilution of pollutants.

Radiocarbon ( $^{14}\text{C}$ ) in EC and OC was measured to distinguish their fossil (mainly coal burning and traffic emissions) and non-fossil sources (mainly biomass burning). The most important contributor to EC was fossil fuel combustion, both in Xi'an and Beijing, contributing  $73 \pm 2\%$  in Xi'an and  $80 \pm 3\%$  in Beijing. The remaining EC arose from biomass burning ( $27 \pm 2\%$  in Xi'an and  $20 \pm 3\%$  in Beijing; Fig. 1). In Xi'an, the fraction of biomass-burning EC in total EC ( $f_{\text{bb}}(\text{EC})$ ) was largely constant during haze and clean periods (range: 25%–29%), regardless of the wide concentration range of EC from biomass burning ( $\text{EC}_{\text{bb}}$ ,  $1.8\text{--}6.4 \mu\text{g m}^{-3}$ ) and fossil fuel combustion ( $\text{EC}_{\text{fossil}}$ ,  $4.3\text{--}18 \mu\text{g m}^{-3}$ ). This suggests that the increase in  $\text{EC}_{\text{fossil}}$  and  $\text{EC}_{\text{bb}}$  concentrations during haze periods in Xi'an is likely caused by the enhanced emissions from both fossil fuel and biomass burning by a similar factor and due to meteorological conditions favoring the accumulation of particulate air pollutions.  $f_{\text{bb}}(\text{EC})$  values in Beijing ( $20 \pm 3\%$  with a range of 17%–24%) were consistently smaller than those in Xi'an (range: 25%–29%), showing that fossil sources contribute more strongly to EC in Beijing. Moreover, during haze periods in Beijing,  $f_{\text{bb}}(\text{EC})$  increased with increasing total EC concentrations (Fig. 2).

In Xi'an, OC concentrations from non-fossil sources averaged  $29 \pm 16 \mu\text{g m}^{-3}$  ( $\text{OC}_{\text{nf}}$ ; range:  $9\text{--}49 \mu\text{g m}^{-3}$ ), slightly higher than those from fossil sources ( $\text{OC}_{\text{fossil}}$ ;  $24 \pm 13 \mu\text{g m}^{-3}$ ; range:  $8\text{--}40 \mu\text{g m}^{-3}$ ) at 95% confidence level (paired  $t$  test,  $p$ -value = 0.01). However, in Beijing,  $\text{OC}_{\text{nf}}$  ( $12 \pm 5 \mu\text{g m}^{-3}$ ;  $3\text{--}19 \mu\text{g m}^{-3}$ ) was significantly lower than  $\text{OC}_{\text{fossil}}$  ( $24 \pm 10 \mu\text{g m}^{-3}$ ;  $4\text{--}33 \mu\text{g m}^{-3}$ ) ( $p$ -value = 0.001). Consequently, the relative contribution of  $\text{OC}_{\text{nf}}$  to total OC ( $f_{\text{nf}}(\text{OC})$ ) was much higher in Xi'an (average  $54 \pm 2\%$ ) than in Beijing ( $34 \pm 3\%$ ).  $f_{\text{nf}}(\text{OC})$  in both cities was considerably higher than the corresponding  $f_{\text{bb}}(\text{EC})$  for all samples (Fig. 1). The main reason for larger  $f_{\text{nf}}(\text{OC})$  than  $f_{\text{bb}}(\text{EC})$  is that primary OC/EC ratios from biomass burning emissions are higher than those from fossil sources. So even though biomass burning contributes a small portion of EC, its contribution to primary OC will be much higher. In addition, other non-combustion sources (e.g., biogenic emissions, cooking fumes) and secondary formation contribute only to OC, but not to EC.

In this study, the  $f_{\text{fossil}}(\text{EC})$  values in Xi'an during winter 2016/2017 are comparable with those previously measured during winter 2015/2016 and winter 2008/2009 (Ni et al., 2018, 2019b), as illustrated in Fig. 3b, pointing to relative constant contribution of fossil fuel combustion vs. biomass burning to EC in Xi'an over the past decade. As shown in Fig. 3b, the  $f_{\text{fossil}}(\text{EC})$  values in Beijing during winter 2016/2017 agree with the values reported at an urban site of Beijing in January 2014 (Fang et al., 2017). A slightly higher  $f_{\text{fossil}}(\text{EC})$  in urban Beijing was observed during February 2010 (Chen et al., 2013). Despite the slight variation of  $f_{\text{fossil}}(\text{EC})$  over time,  $f_{\text{fossil}}(\text{EC})$  in Beijing is generally higher than that in Xi'an (Fig. 3b). The presented overall average  $f_{\text{fossil}}(\text{OC})$  for winter 2016/2017 in Beijing ( $66 \pm 3\%$ ) was higher than that in Xi'an ( $46 \pm 3\%$ ), consistent with previously reported  $f_{\text{fossil}}(\text{OC})$  in Beijing and Xi'an (Zhang et al., 2015; Ni et al., 2019a). Lower  $f_{\text{fossil}}(\text{OC})$  values in winter were reported for Chongqing (24%), and higher  $f_{\text{fossil}}(\text{OC})$  was observed in Taiyuan (71%) during winter 2013/2014 (Ni et al., 2019a). The comparison of  $f_{\text{fossil}}(\text{EC})$  and  $f_{\text{fossil}}(\text{OC})$  in different Chinese cities indicates that the relative importance of fossil sources in carbonaceous aerosols vary-varies spatially, and can change over the years. In Xi'an, clean periods showed a slight day-night difference with increased contributions of fossil sources to EC and OC during the day. During haze periods, especially the 2<sup>nd</sup> haze event (XH\_day2, XH\_night2), this day-night difference disappeared, which suggests a long residence

215 time of the pollution particles in the urban atmosphere during haze events.

Overall, our  $^{14}\text{C}$  data show that fossil sources contribute more strongly to EC and OC in Beijing than in Xi'an, which is consistent with previous observations. Both in Beijing and Xi'an the fossil vs. non-fossil contributions to EC and OC did not change drastically during haze and clean periods. In Xi'an, a slight day-night difference was observed during clean periods, but disappeared during haze periods, suggesting a large accumulation of particles.

### 220 3.2 Fossil EC apportioned by stable carbon isotopes: coal vs. liquid fossil fuel

Besides  $\text{F}^{14}\text{C}$ , the  $\delta^{13}\text{C}$  of EC adds additional dimension where fossil EC can be distinguished into EC from burning of coal and of liquid fossil fuel (i.e., vehicle emissions). Considerable geographical differences in  $\delta^{13}\text{C}_{\text{EC}}$  signatures were observed, with more depleted values in Xi'an ( $-25.1 \pm 0.5\text{‰}$ ;  $-25.6\text{‰}$  to  $-24.4\text{‰}$ ) relative to those in Beijing ( $-24.1 \pm 0.4\text{‰}$ ;  $-24.4\text{‰}$  to  $-23.4\text{‰}$ ; Fig. 3). The Xi'an signatures are consistent with the signature of liquid fossil fuel combustion ( $\delta^{13}\text{C}_{\text{liq.fossil}} = -25.5$   
225  $\pm 1.3\text{‰}$ ; Sect. 2.4), whereas the more enriched values in Beijing indicate the influence of coal combustion ( $\delta^{13}\text{C}_{\text{coal}} = -23.4 \pm 1.3\text{‰}$ ).

In both Xi'an and Beijing, moderate differences exist in  $\delta^{13}\text{C}_{\text{EC}}$  between clean and haze days, pointing to a shift in combustion sources. In Xi'an,  $\delta^{13}\text{C}_{\text{EC}}$  during clean periods ( $\sim -25.5\text{‰}$ ) was slightly depleted compared to that during haze periods ( $-25.0\text{‰}$  to  $-24.4\text{‰}$ ), whereas Beijing exhibited more enriched  $\delta^{13}\text{C}_{\text{EC}}$  during clean periods ( $-23.4\text{‰}$ ) than during haze periods ( $-24.4\text{‰}$   
230 to  $-24.1\text{‰}$ ). This suggests a moderate increase in coal combustion contribution to EC in Xi'an during haze days and a decrease in Beijing. In Xi'an, no strong day-night difference in  $\delta^{13}\text{C}_{\text{EC}}$  was observed, with the largest absolute differences of  $0.5\text{‰}$  between XH\_day1 and XH\_night1. The day-night differences are small relative to the uncertainties of the potential sources; for example, the endmember range for coal combustion is more uncertain ( $\pm 1.3\text{‰}$ ). The small day-night differences in  $\delta^{13}\text{C}_{\text{EC}}$  reflect well-mixed EC emissions.

235 The Bayesian MCMC model takes into account the uncertainties of the  $\delta^{13}\text{C}$  and  $\text{F}^{14}\text{C}$  endmembers and statistically apportions EC into the fraction of biomass burning ( $f_{\text{bb}}$ ), coal combustion ( $f_{\text{coal}}$ ) and liquid fossil fuel combustion ( $f_{\text{liq.fossil}}$ ). The MCMC-derived  $f_{\text{bb}}$  is in principle the same as the  $^{14}\text{C}$ -based  $f_{\text{bb}}(\text{EC})$  (Fig. S3). The MCMC results (Fig. 4) show that there were no strong day-night differences in EC sources during haze and clean periods in Xi'an. Liquid fossil fuel combustion was the most important contributor to EC in Xi'an, with increased contribution during clean periods. In Beijing, coal combustion was the  
240 dominant source of EC, with the relative contribution ranging from 48% (median; 31%–61%, interquartile range) during haze periods to 61% (45%–71%) during clean periods.  $f_{\text{bb}}$  was fairly constant between haze and clean periods with respect to  $f_{\text{coal}}$  and  $f_{\text{liq.fossil}}$  for all samples. In Xi'an,  $f_{\text{bb}}$  was comparable to  $f_{\text{coal}}$  during haze days, and larger than  $f_{\text{coal}}$  during clean days. In Beijing, biomass-burning EC was the smallest fraction in total EC, with smaller  $f_{\text{bb}}$  than  $f_{\text{coal}}$  during both haze and clean days. Concentrations of total EC increased by 2 times from clean days ( $\sim 7.4 \mu\text{g m}^{-3}$ ) to haze days ( $18.0 \mu\text{g m}^{-3}$ ) in Xian, and 8 times  
245 in Beijing ( $1.6 \mu\text{g m}^{-3}$  to  $13.5 \mu\text{g m}^{-3}$ ). The increased EC concentrations during haze periods in Xi'an were attributed to liquid fossil fuel combustion (43%), coal combustion (29%) and biomass burning (28%). However, in Beijing, coal combustion

contributed most of the increased concentrations of EC (45%), followed by burning of liquid fossil fuel (33%) and biomass (22%).

In summary, complementing  $^{14}\text{C}$  with  $\delta^{13}\text{C}$  allows for quantitative constraints on EC sources: EC was dominated by liquid fossil fuel combustion (i.e., vehicle emissions) in Xi'an and by coal burning in Beijing, especially during clean periods. In Xi'an, no strong day-night differences in EC sources were observed during haze and clean periods. Compared with earlier observations in Xi'an (Fig. 3b), we found that the  $\delta^{13}\text{C}_{\text{EC}}$  values in January 2017 from this study are comparable with wintertime  $\delta^{13}\text{C}_{\text{EC}}$  in 2015/2016 (Ni et al., 2019b), but much more depleted than wintertime  $\delta^{13}\text{C}_{\text{EC}}$  in 2008/2009 (Ni et al., 2018) and January 2003 (Cao et al., 2011). This suggests that fossil sources of EC in Xi'an have changed in the past decade, with decreasing relative contribution from coal combustion. This is in line with recent changes in energy use, and the decreasing enrichment factors of As and Pb (e.g., i.e., indicators of coal combustion) in Xi'an, as documented in recent studies (Xu et al., 2016). As shown in Fig. 3b, in Beijing, variations in  $\delta^{13}\text{C}_{\text{EC}}$  from January 2003 (Cao et al., 2011) to January 2017 (this study) are much narrower than those in Xi'an, indicating that EC combustion sources did not change significantly throughout the years in Beijing. Our  $\delta^{13}\text{C}_{\text{EC}}$  values overlap with those in January 2014 (Fang et al., 2017) and fall into the range of reported  $\delta^{13}\text{C}_{\text{EC}}$  values in urban Beijing (Cao et al., 2011; Chen et al., 2013) and the regional receptor site of Beijing (Andersson et al., 2015; Fang et al., 2017).

### 3.3 Primary and secondary OC

As explained in Sect. 2.4,  $\text{OC}_{\text{nf}}$  and  $\text{OC}_{\text{fossil}}$  are apportioned into primary ( $\text{POC}_{\text{bb}}$ ,  $\text{POC}_{\text{fossil}}$ ) and secondary OC ( $\text{OC}_{\text{o,nf}}$ ,  $\text{SOC}_{\text{fossil}}$ ; Fig. 5). The large error bars of their concentrations reflect the large uncertainties in  $r_{\text{bb}}$  and  $r_{\text{fossil}}$ . It should be noticed that  $\text{OC}_{\text{o,nf}}$  is used as an approximation of  $\text{SOC}_{\text{nf}}$ , or can be regarded as an upper limit of  $\text{SOC}_{\text{nf}}$  if cooking is a prominent OC source.

In Xi'an, both ratios of  $\text{OC}_{\text{o,nf}}/\text{POC}_{\text{bb}}$  and  $\text{SOC}_{\text{fossil}}/\text{POC}_{\text{fossil}}$  increased during haze periods (Fig. 5a).  $\text{OC}_{\text{o,nf}}/\text{POC}_{\text{bb}}$  ratio increased by 2.5 times from 0.33–0.46 during clean periods to 0.86–1.1 during haze periods; in contrast  $\text{SOC}_{\text{fossil}}/\text{POC}_{\text{fossil}}$  increased by 1.5 times from 0.46–0.50 to 0.62–0.78. This underlines that haze episodes in Xi'an were mainly caused by additional SOC formation, with larger contribution from non-fossil sources than fossil sources. As shown in Fig. 5b, the contribution of SOC (i.e.,  $\text{SOC} \cong \text{OC}_{\text{o,nf}} + \text{SOC}_{\text{fossil}}$ ) to OC increased from clean periods (28%–32%) to haze periods (44%–48%); this mainly resulted from increased contribution of  $\text{OC}_{\text{o,nf}}$  to total OC (i.e., from 14%–16% to 26%–29%). In Xi'an, the day-night difference was larger during clean periods with less SOC at night for both absolute concentration and relative contribution to total OC (Figs. 5b, 5c).

In contrast, Beijing had the opposite variation trends of  $\text{OC}_{\text{o,nf}}/\text{POC}_{\text{bb}}$  and  $\text{SOC}_{\text{fossil}}/\text{POC}_{\text{fossil}}$  from clean to haze periods.  $\text{OC}_{\text{o,nf}}/\text{POC}_{\text{bb}}$  ratios during clean periods (1.3) were on average five times higher than those during haze periods (0.18–0.33), and  $\text{SOC}_{\text{fossil}}/\text{POC}_{\text{fossil}}$  ratios during clean periods (0.71) were slightly higher than those during haze periods (0.41–0.64). This suggests that in Beijing the increased OC concentrations during haze periods were mainly derived from elevated concentrations of  $\text{POC}_{\text{bb}}$  and  $\text{POC}_{\text{fossil}}$ . As shown in Fig. 5b, high SOC contribution to total OC was observed during clean periods, mainly

due to elevated contribution from  $OC_{o,nf}$ . The  $OC_{o,nf}$  is not likely attributed to biogenic OC, because the biogenic emissions are very low in winter. As a result, the elevated contribution from  $OC_{o,nf}$  to OC during clean periods in Beijing could be attributed to regional sources. During clean periods, concentrations of OC and  $OC_{o,nf}$  are small, and the measured carbon concentrations can reflect regional sources, which are dominated by secondary sources due to long-range transport. It could also be that contribution of cooking OC to  $OC_{o,nf}$  can be noticeable during clean conditions.

The fossil fraction of the total SOC can be defined as  $f_{fossil}(SOC) = SOC_{fossil}/SOC$ . In Xi'an around half of SOC was derived from fossil sources ( $f_{fossil}(SOC) = 44 \pm 6\%$ ), whereas  $f_{fossil}(SOC) = 75 \pm 10\%$  in Beijing. Using a similar approach with this study, Zhang et al. (2015) also found that Beijing had higher  $f_{fossil}(SOC)$  (48%–63%) than in Xi'an (30%–35%) during the January 2013 severe haze events. These findings suggest the important contribution of fossil sources to SOC in Beijing and non-fossil sources in Xi'an.  $f_{fossil}(SOC)$  in Beijing increased during haze periods, whereas the opposite trend was found in Xi'an (Fig. 6). During haze periods in Beijing,  $f_{fossil}(SOC)$  overlapped with  $f_{fossil}(EC)$ , and was clearly higher than  $f_{fossil}(OC)$ .

Together these results reveal the differences in primary and secondary OC in two Chinese megacities. The contribution of SOC to total OC increased from clean to haze periods in Xi'an. In contrary, SOC/OC ratios increased from haze to clean periods in Beijing, mainly due to increased SOC from non-fossil sources. SOC was dominated by fossil sources in Beijing but by non-fossil sources in Xi'an, especially during haze periods. In Xi'an, the day-night difference was larger during clean periods with less SOC at night.

### 3.4 Differences between the fractions of non-fossil carbon in OC and EC

The differences between  $f_{nf}(OC)$  and  $f_{bb}(EC)$  were smaller in Beijing, ranging from 11% to 20%, compared to 25%–29% in Xi'an. To better understand what governs the differences, we express  $f_{nf}(OC)$  in terms of fossil to biomass burning ratio in EC and primary OC/EC emissions ratios. Starting from the formulas of  $f_{bb}(EC)$  and  $f_{nf}(OC)$ :

$$f_{bb}(EC) = \frac{EC_{bb}}{EC_{bb} + EC_{fossil}} = \frac{1}{1 + \frac{EC_{fossil}}{EC_{bb}}} \quad (14)$$

$$f_{nf}(OC) = \frac{OC_{nf}}{OC_{nf} + OC_{fossil}} = \frac{POC_{bb} + OC_{o,nf}}{POC_{bb} + OC_{o,nf} + POC_{fossil} + SOC_{fossil}} \quad (15)$$

We find that

$$f_{nf}(OC) = \frac{1}{1 + \frac{(1 + SOC_{fossil}/POC_{fossil})}{(1 + OC_{o,nf}/POC_{bb})} \times \frac{POC_{fossil}}{POC_{bb}}} = \frac{1}{1 + \frac{(1 + SOC_{fossil}/POC_{fossil})}{(1 + OC_{o,nf}/POC_{bb})} \times \frac{r_{fossil}}{r_{bb}} \times \frac{EC_{fossil}}{EC_{bb}}} \quad (16)$$

where  $r_{fossil}$  is the weighted average of  $r_{coal}$  and  $r_{vehicle}$ .

Comparing Eq. (14) with Eq. (16), we find that  $f_{nf}(OC)$  and  $f_{bb}(EC)$  would be equal if  $\frac{(1+SOC_{fossil}/POC_{fossil})}{(1+OC_{o,nf}/POC_{bb})} \times \frac{r_{fossil}}{r_{bb}} = 1$ . Since  $r_{fossil}$  is usually smaller than  $r_{bb}$ ,  $f_{nf}(OC)$  tends to be larger than  $f_{bb}(EC)$ , assuming that SOC formation is comparable for fossil or non-fossil sources (i.e.,  $\frac{(1+SOC_{fossil}/POC_{fossil})}{(1+OC_{o,nf}/POC_{bb})} \sim 1$ ). With smaller  $r_{fossil}$  than  $r_{bb}$ , similar  $f_{nf}(OC)$  and  $f_{bb}(EC)$  can result from larger secondary formation from fossil sources than non-fossil sources (i.e.,  $\frac{(1+SOC_{fossil}/POC_{fossil})}{(1+OC_{o,nf}/POC_{bb})} > 1$ ). However, the fossil source coal combustion has a higher primary OC to EC ratio than vehicle emissions (i.e.,  $r_{coal} > r_{vehicle}$ ). Therefore, in a city where biomass burning and coal combustion are the dominant pollution sources,  $f_{nf}(OC)$  and  $f_{bb}(EC)$  will be more similar than in a city where the main sources are biomass burning and vehicle emissions.

Compared to Xi'an, Beijing had significantly smaller differences between  $f_{bb}(EC)$  and  $f_{nf}(OC)$  (Fig. 1), which was also observed in previous studies during the haze event in January 2013 (Zhang et al., 2015). Comparing Eq. (14) with Eq. (16), this suggests either strong contribution from coal combustion in Beijing or large secondary formation from fossil sources, or both. The stronger contribution of coal combustion to OC in Beijing than in Xi'an was a direct consequence of a larger proportion of coal combustion in EC in Beijing, as demonstrated by the Bayesian MCMC results of EC (Sect. 3.2). The latter was further validated by the variation of SOC. The  $f_{fossil}(SOC)$  in Beijing was higher than that in Xi'an, despite the variations between haze and clean periods (Sect. 3.3). By combining  $^{14}C$  measurements with other state-of-art analytical techniques (e.g., aerosol mass spectrometry), Huang et al. (2014) also found that fossil OC was mostly secondary in nature in Beijing, and non-fossil SOC formation was dominant in Xi'an during a wintertime haze episode (i.e., Beijing had a larger  $f_{fossil}(SOC)$  than Xi'an). However, atmospheric mechanisms responsible for the enhancement in fossil-derived [SOA—secondary organic aerosol](#) formation in Beijing remain unclear.

Furthermore, as shown in Fig. 1, unlike Xi'an where the differences between  $f_{nf}(OC)$  and  $f_{bb}(EC)$  were relatively constant for all samples, in Beijing the differences between  $f_{nf}(OC)$  and  $f_{bb}(EC)$  were smaller during haze periods than clean periods, caused by decreased  $f_{bb}(EC)$  and slightly increased  $f_{nf}(OC)$  during clean periods. This might indicate a higher relative contribution from coal combustion and/or fossil-dominated SOC during haze periods in Beijing. However, the Bayesian MCMC results of EC show the opposite, i.e., in Beijing the contribution of coal combustion to EC was lower during haze periods than during clean periods (Sect. 3.2). Therefore, the only possible explanation is that, during haze periods in Beijing, SOC was dominated by fossil sources. This is validated by significantly larger  $f_{fossil}(SOC)$  during haze periods (76%–81%) than during clean periods (~55%; Sect. 3.3).

In conclusion, this section discusses the factors that governs the differences between  $f_{bb}(EC)$  and  $f_{nf}(OC)$ , and concludes that smaller differences suggest stronger contribution from coal combustion and/or larger secondary formation from fossil sources. This is further examined and validated by source apportionment results of EC and OC (Sect. 3.1–3.3).

## 4 Conclusion

In this study the sources of carbonaceous aerosol were quantified using a dual-carbon isotopic approach for PM<sub>2.5</sub> samples collected in urban Xi'an and Beijing reaching "red alarm" level during December 2016 and January 2017. The <sup>14</sup>C results showed that fossil sources dominated EC, contributing on average  $73 \pm 2\%$  of EC in Xi'an and  $80 \pm 3\%$  of EC in Beijing. The remaining EC was attributed to biomass burning. In Xi'an,  $f_{bb}(EC)$  was fairly constant during haze and clean periods, despite the wide range of EC concentrations. However, in Beijing,  $f_{bb}(EC)$  increased with increasing EC concentrations. Complementing <sup>14</sup>C with  $\delta^{13}C$  in a Bayesian MCMC approach allows for separation of fossil sources of EC into coal combustion and liquid fossil fuel combustion. The MCMC results in Xi'an suggest that liquid fossil fuel combustion contributed 44%–49% of EC during haze periods, and 54%–57% of EC during clean periods. In Beijing, coal combustion was the ~~dominate~~-dominating fossil source of EC, with decreasing contribution to EC from clean periods (~61%) to haze periods (~48%).

<sup>14</sup>C measurements of OC showed that the contribution of non-fossil sources to OC was larger than that to EC, and was on average  $54 \pm 2\%$  in Xi'an and  $34 \pm 3\%$  in Beijing. The differences between non-fossil fraction in OC and EC were smaller in Beijing and larger in Xi'an. In Xi'an, the fraction of SOC in total OC was larger during haze periods than during clean periods, mainly due to increased SOC from non-fossil sources. Beijing showed the opposite trends with a larger fraction of SOC in total OC during clean periods than during haze periods, mainly due to elevated contribution from non-fossil SOC during clean periods.

SOC was dominated by non-fossil sources in Xi'an but by fossil sources in Beijing, especially during haze periods. The relative contribution of fossil sources to SOC ( $f_{fossil}(SOC)$ ) was consistently higher in Beijing than in Xi'an. In Beijing,  $f_{fossil}(SOC)$  was higher during haze periods (76%–81%) than during clean periods (55%), whereas an opposite trend was found in Xi'an, with  $f_{fossil}(SOC)$  increasing from ~39%–43% during haze periods to ~52% during clean periods. In Xi'an, a slight day-night difference was found during clean periods, with increasing fossil contribution to OC and EC during the day and less SOC at night. During strong haze, this day-night difference was negligible, suggesting a large accumulation under stagnant weather conditions during the severe haze periods.

*Data availability.* Data used to support the findings in this study are archived at the Institute of Earth Environment, Chinese Academy of Sciences, and are available on request from the corresponding author.

*Competing interests.* The authors declare that they have no conflict of interest.

360 *Author contributions.* RJH and UD designed the study. Isotope measurements were made by HN, MMC, and JG. Data analysis and interpretation were made by HN, MMC, RJH, and UD. HN wrote the paper with contributions from all co-authors.

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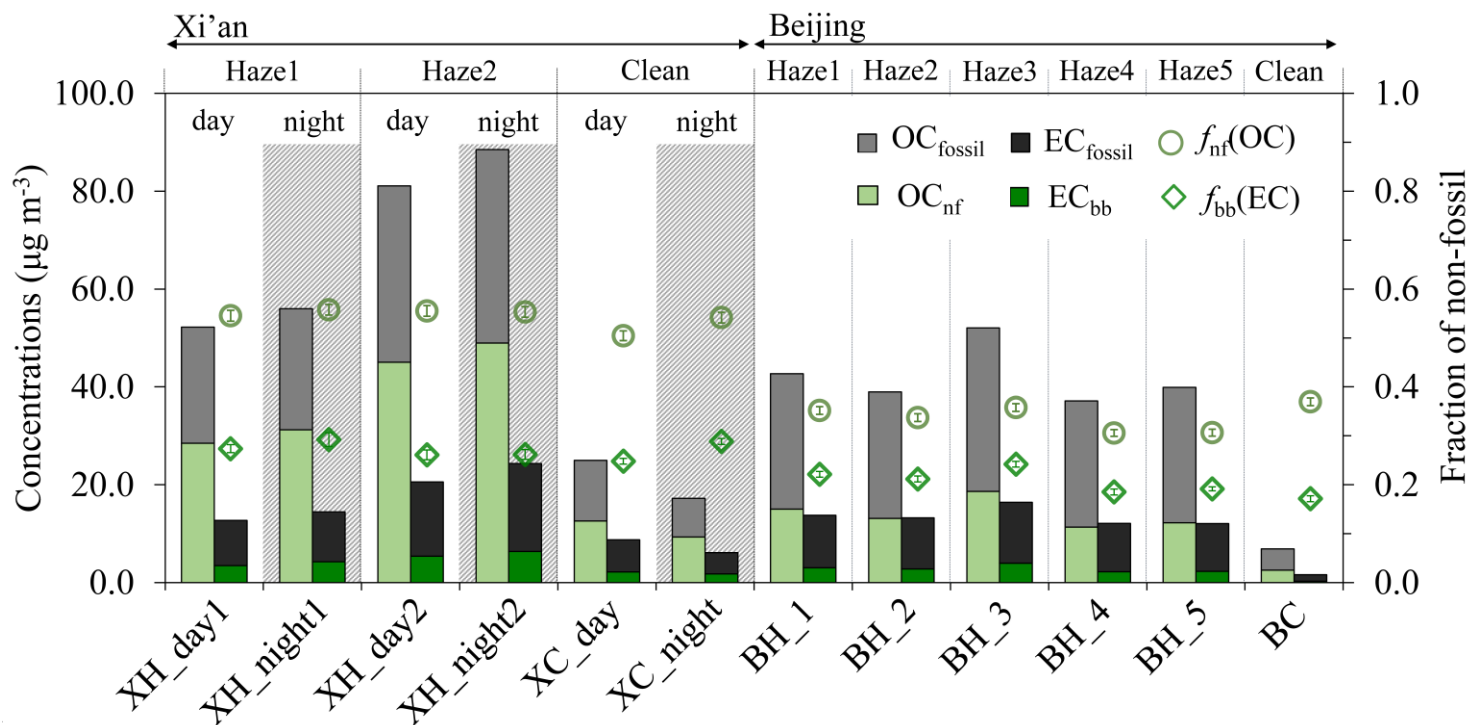
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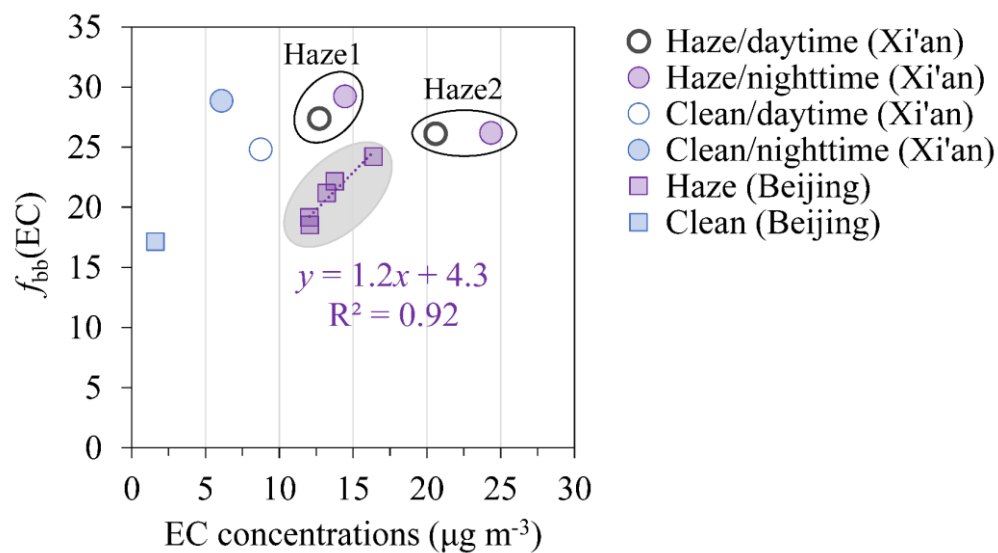
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**Table 1.** Equations for  $^{14}\text{C}$  source apportionment of EC and OC. See Sect. 2.4 for details.  $r_{\text{bb}}$  and  $r_{\text{fossil}}$  are primary OC/EC ratio for biomass burning and fossil fuel combustion, respectively. Estimation of  $r_{\text{bb}}$  and  $r_{\text{fossil}}$  is presented in Supplement S2.

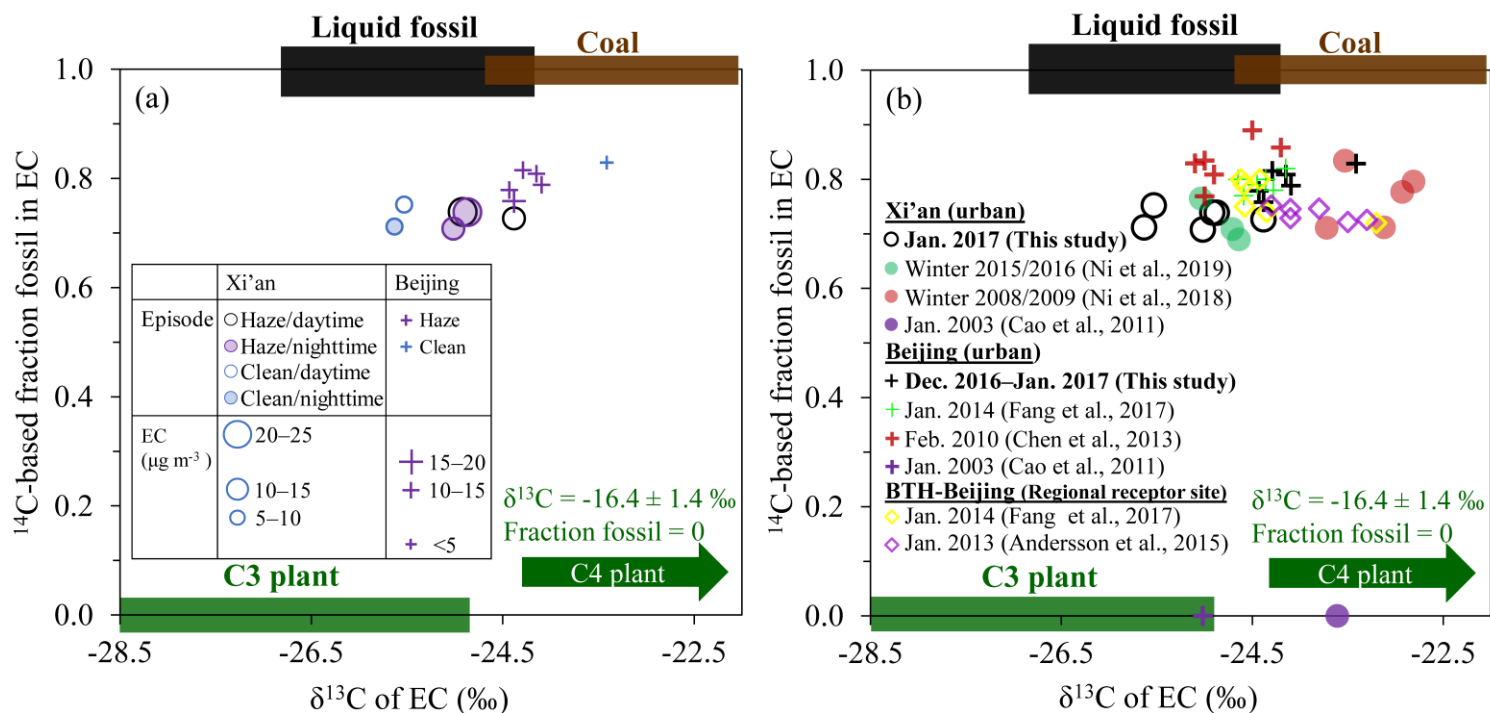
Equations	
$\text{EC}_{\text{bb}} = \text{EC} \times f_{\text{bb}}(\text{EC})$	(3)
$\text{EC}_{\text{fossil}} = \text{EC} \times (1 - f_{\text{bb}}(\text{EC})) = \text{EC} \times f_{\text{fossil}}(\text{EC})$	(4)
$\text{OC}_{\text{nf}} = \text{OC} \times f_{\text{nf}}(\text{OC})$	(5)
$\text{OC}_{\text{fossil}} = \text{OC} \times (1 - f_{\text{nf}}(\text{OC})) = \text{OC} \times f_{\text{fossil}}(\text{OC})$	(6)
$\text{POC}_{\text{bb}} = \text{EC}_{\text{bb}} \times r_{\text{bb}}$	(7)
$\text{OC}_{\text{o,nf}} = \text{OC}_{\text{nf}} - \text{POC}_{\text{bb}}$	(8)
$\text{POC}_{\text{fossil}} = \text{EC}_{\text{fossil}} \times r_{\text{fossil}}$	(9)
$\text{SOC}_{\text{fossil}} = \text{OC}_{\text{fossil}} - \text{POC}_{\text{fossil}}$	(10)
$\text{SOC} = \text{SOC}_{\text{nf}} + \text{SOC}_{\text{fossil}} \cong \text{OC}_{\text{o,nf}} + \text{SOC}_{\text{fossil}}$	(11)
$f_{\text{fossil}}(\text{SOC}) = \text{SOC}_{\text{fossil}} / \text{SOC}$	(12)



**Figure 1.** Mass concentrations of OC and EC from fossil and non-fossil sources (OC<sub>fossil</sub>, OC<sub>nf</sub>, EC<sub>fossil</sub> and EC<sub>bb</sub>) as well as fraction of non-fossil carbon in OC and EC ( $f_{\text{nf}}(\text{OC})$  and  $f_{\text{bb}}(\text{EC})$ , respectively) for daytime and nighttime PM<sub>2.5</sub> samples in Xi'an, and 24h-integrated PM<sub>2.5</sub> samples in Beijing during haze and clean periods during the measurement periods (2 December 2016 to 10 January 2017). For each city “haze” and “clean” are used to represent high and low pollution events, and clean days at each site are defined as days with PM<sub>2.5</sub> < median concentration in the winter heating season from 15 November 2016 to 15 March 2017. Uncertainties of <sup>14</sup>C-apportioned  $f_{\text{nf}}(\text{OC})$  and  $f_{\text{bb}}(\text{EC})$  are indicated but are too small to be visible. The data are shown in Table S3.

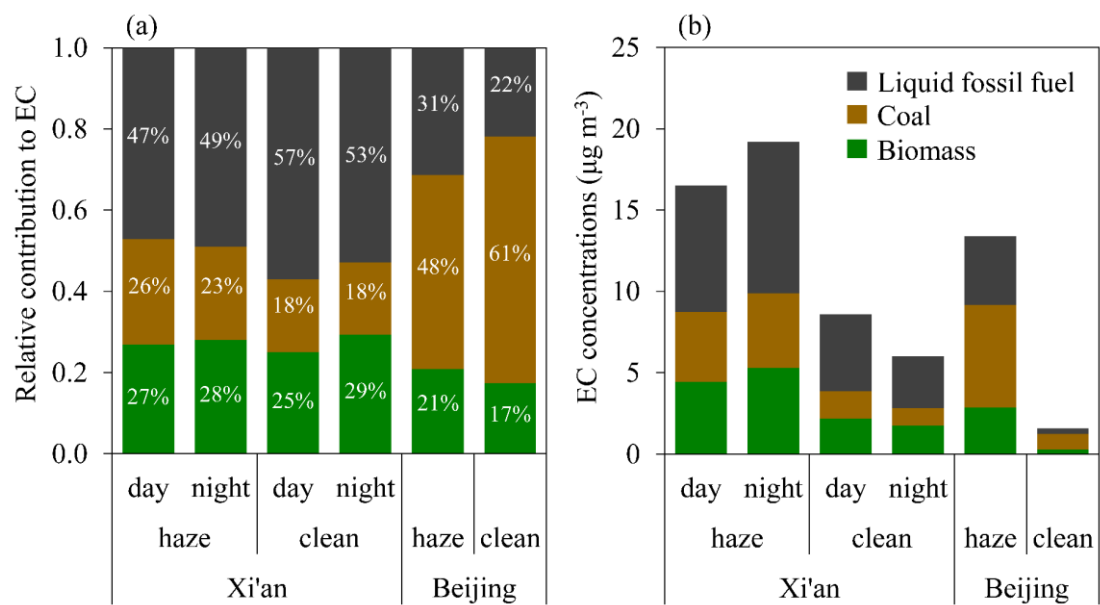


540 **Figure 2.** Scatter plot of the relative contribution of biomass burning to EC ( $f_{bb}(EC)$ ; %) against EC concentrations.

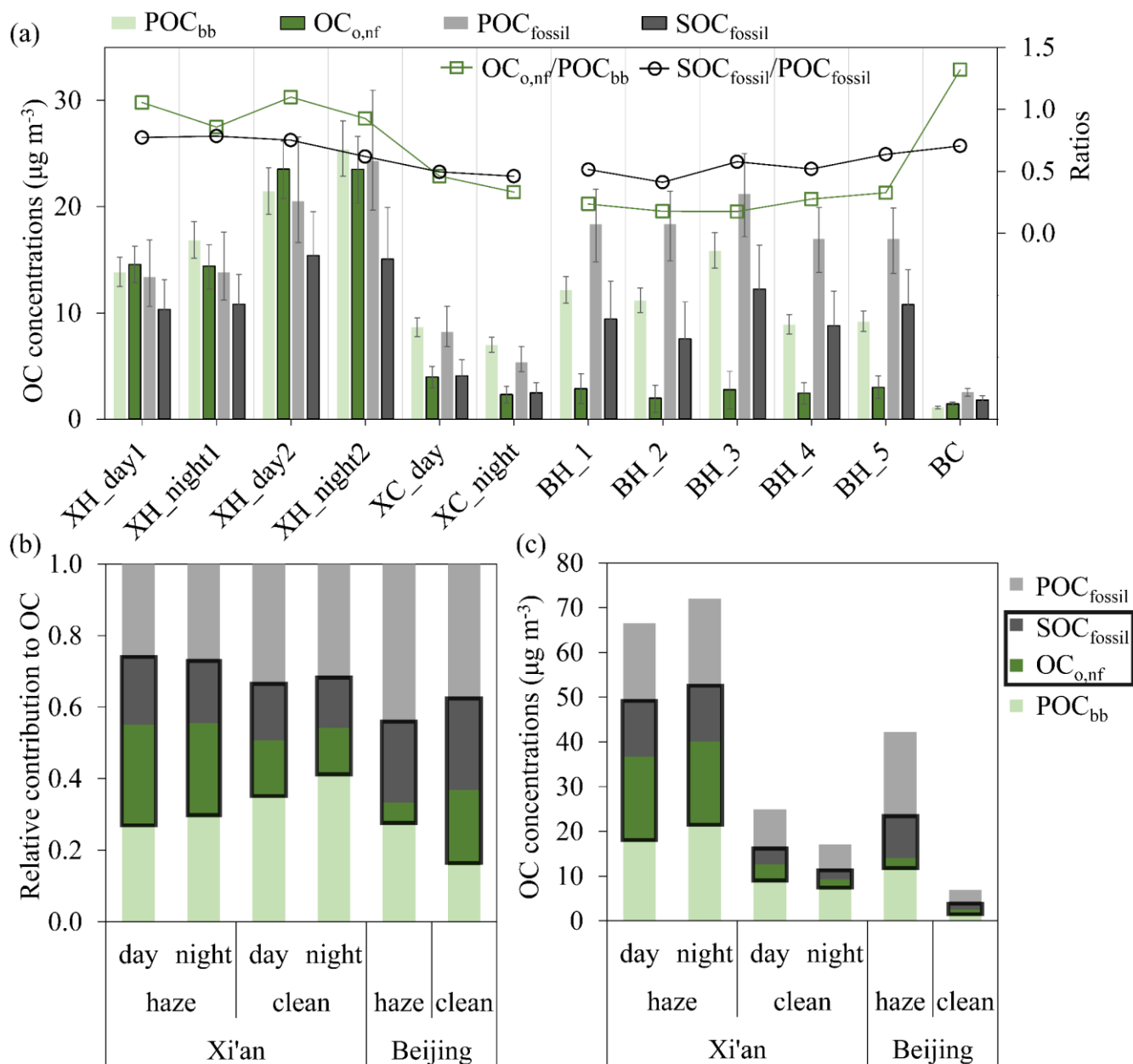


**Figure 3.** (a)  $^{14}\text{C}$ -based fraction fossil versus  $\delta^{13}\text{C}$  for EC during haze and clean periods in Xi'an and Beijing, China. The symbol size is an indicator of EC concentrations. (b) Comparison with previous observations in Xi'an and Beijing, where BTH-Beijing is a regional receptor site of Beijing, located at 100 km southwest of Beijing. Samples from Cao et al. (2011) are placed on the  $x$ -axis, because no  $^{14}\text{C}$  data were available. The expected  $^{14}\text{C}$  and  $\delta^{13}\text{C}$  endmember ranges for emissions from C3 plant burning, liquid fossil fuel burning and coal burning are shown as green, black and brown bars, respectively. The  $\delta^{13}\text{C}$  source signatures are indicated as mean  $\pm$  SD (Sect. 2.4). The  $\delta^{13}\text{C}$  signature of corn stalk burning (i.e., C4 plant;  $-16.4 \pm 1.4 \text{ ‰}$ ) is also indicated.

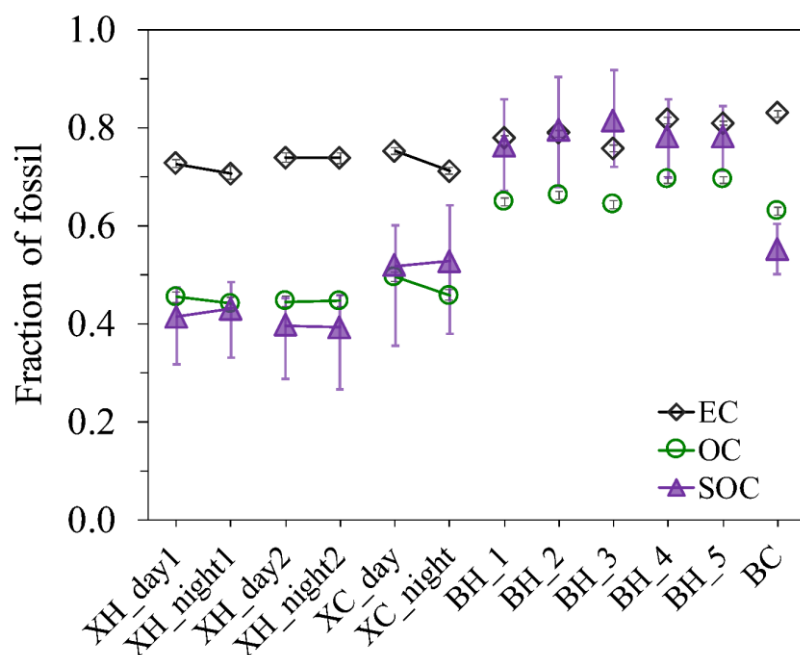




550 **Figure 4.** (a) Fractional contributions of three combustion sources to EC during haze and clean periods in Xi'an and Beijing.  
 (b) EC concentrations ( $\mu\text{g m}^{-3}$ ) from each combustion source. The data are presented in Tables S5 and S6.



**Figure 5.** (a) Concentrations of  $\text{POC}_{\text{bb}}$ ,  $\text{OC}_{\text{o,nf}}$ ,  $\text{POC}_{\text{fossil}}$  and  $\text{SOC}_{\text{fossil}}$  ( $\mu\text{g m}^{-3}$ ), and the mass ratio of  $\text{OC}_{\text{o,nf}}/\text{POC}_{\text{bb}}$  and  $\text{SOC}_{\text{fossil}}/\text{POC}_{\text{fossil}}$  of each sample. The error bars indicate the interquartile range (25th–75th percentile) of the median concentrations. Averaged fraction (b) and concentration (c) of  $\text{POC}_{\text{bb}}$ ,  $\text{OC}_{\text{o,nf}}$ ,  $\text{POC}_{\text{fossil}}$  and  $\text{SOC}_{\text{fossil}}$  in total OC during haze and clean periods in Xi'an and Beijing, China. The data are given in Table S4.



**Figure 6.** Fraction of fossil carbon in EC, OC and SOC ( $f_{\text{fossil}}(\text{EC})$ ,  $f_{\text{fossil}}(\text{OC})$  and  $f_{\text{fossil}}(\text{SOC})$ , respectively) during haze and clean periods in Xi'an and Beijing. Interquartile ranges (25th–75th percentile) of the median  $f_{\text{fossil}}(\text{SOC})$  are shown as vertical bars in purple. Uncertainties of  $^{14}\text{C}$ -apportioned  $f_{\text{fossil}}(\text{EC})$  and  $f_{\text{fossil}}(\text{OC})$  are indicated but are too small to be visible. The studied samples include day and night samples in Xi'an (X) and 24 h integrated samples in Beijing (B) during haze periods (“H” samples) and clean periods (“C” samples). For details of each sample (x-axis), see Fig. 1 and Table S1.

*Supplement of*

**Measurement report: Dual-carbon isotopic characterization of carbonaceous aerosol reveals [s](#) different primary and secondary sources in Beijing and Xi'an during severe haze events**

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## S1. Evaluation of potential interference from carbonates

PM<sub>2.5</sub> samples with the highest concentrations of calcium (associated with carbonate in most cases; Chow and Watson, 2002) in Xi'an, China were used to check the presence of carbonate carbon. Carbonate carbon in an aerosol sample can be verified if there is significant difference in TC mass before and after acidification (i.e., expose the sample to HCl vapor; NIOSH, 1999). Negligible carbonate carbon was found, because TC mass before and after acidification agrees with each other within the measurement uncertainty. PM<sub>2.5</sub> samples in Beijing are less likely affected by carbonates, owing to lower dust concentrations (i.e., a major source of carbonate) than in Xi'an (Huang et al., 2014). Therefore, acidification to remove interferences from carbonates is not necessary, because carbonate carbon is very small compared to the relatively larger OC and EC amounts for both mass determination and carbon isotopic analysis.

## S2. Estimation of $r_{bb}$ and $r_{fossil}$

$r_{bb}$  and  $r_{fossil}$  are primary OC/EC ratio for biomass burning and fossil fuel combustion, respectively.  $r_{fossil}$  is the weighted average of (POC/EC)<sub>coal</sub> (i.e.,  $r_{coal}$ ), and (POC/EC)<sub>vehicle</sub> ( $r_{vehicle}$ ). The weight associated with  $r_{coal}$  (denoted as  $w_{coal}$ ) is the relative contribution of coal combustion to fossil EC. That is,

$$w_{coal} = \frac{EC_{coal}}{EC_{fossil}} = \frac{EC_{coal}}{EC_{coal} + EC_{liq.fossil}} \quad (S1)$$

where  $EC_{fossil}$  is the sum of  $EC_{coal}$  and EC from liquid fossil fuel combustion (i.e., vehicle emissions;  $EC_{liq.fossil}$ ).

Eq. (S1) can be formulated as:

$$w_{coal} = \frac{f_{coal}}{f_{fossil}} = \frac{f_{coal}}{f_{coal} + f_{liq.fossil}} \quad (S2)$$

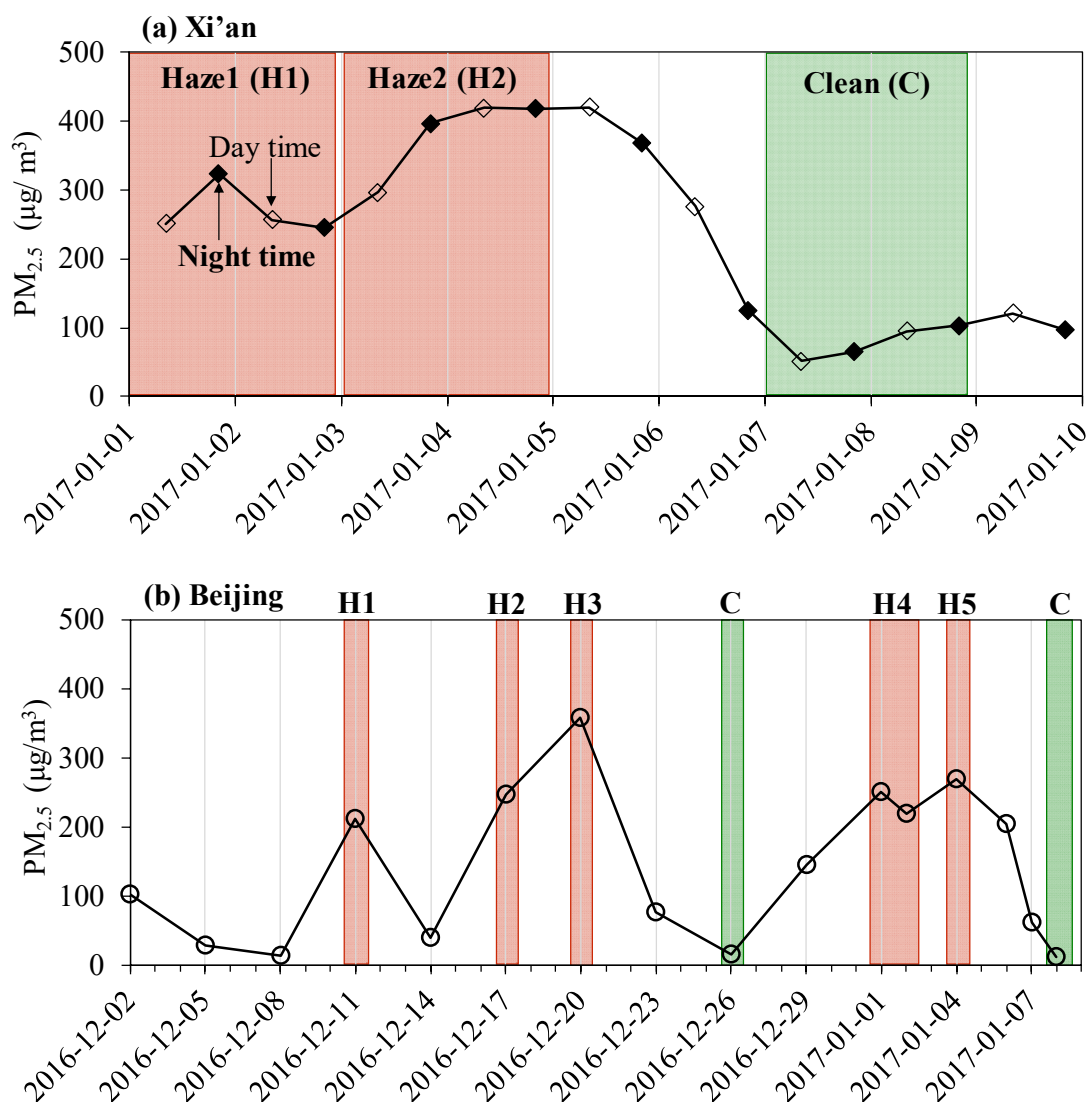
where  $f_{coal}$  and  $f_{liq.fossil}$  are the relative contribution of coal combustion and liquid fossil fuel combustion to EC, respectively. The sum of  $f_{coal}$  and  $f_{liq.fossil}$  is  $f_{fossil}$  of EC, which is well constrained by F<sup>14</sup>C of EC. The [probability density functions \(PDFs\)](#) of  $f_{coal}$  and  $f_{liq.fossil}$  (e.g., Fig. S4), derived from the Bayesian calculations detailed in Sect. 2.4 of the main text, are used to calculate the PDFs of  $w_{coal}$ .

Best estimates of  $r_{bb}$  ( $4 \pm 1$ ; average  $\pm$  SD),  $r_{coal}$  ( $2.38 \pm 0.44$ ), and  $r_{vehicle}$  ( $0.85 \pm 0.16$ ) are defined through a literature search as described in Ni et al. (2018), and their values are comparable to values used in earlier studies (Zhang et al., 2014, 2015).

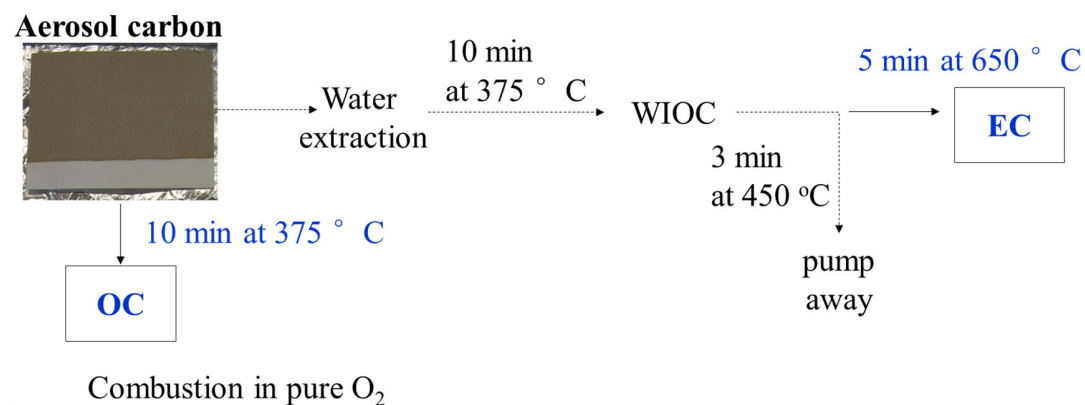
### S3. Uncertainties of $^{14}\text{C}$ source apportionment results

To propagate uncertainties, we conducted a Monte Carlo simulation with 10000 individual calculations to propagate experimental uncertainties and uncertainties in parameters (e.g.,  $F^{14}\text{C}_{\text{nf}}$ ,  $r_{\text{bb}}$  and  $r_{\text{fossil}}$ ) following Eq. (3–12).

For each calculation,  $F^{14}\text{C}$  and mass of EC and OC were randomly chosen from a normal distribution symmetric around the measured values with the experimental uncertainties as standard deviation (SD). For  $F^{14}\text{C}_{\text{nf}}$ ,  $r_{\text{bb}}$ ,  $r_{\text{coal}}$  and  $r_{\text{vehicle}}$ , random values of each parameter were chosen from a triangular frequency distribution with its maximum frequency at the central value and 0 frequency at the lower limit and upper limit of each parameter. For  $w_{\text{coal}}$ , random values from the respective [probability density function \(PDFs\)](#) of  $w_{\text{coal}}$  were used (Supplement S2). In this way 10000 random sets of variables were generated. For  $f_{\text{bb}}(\text{EC})$ ,  $f_{\text{nf}}(\text{OC})$ ,  $\text{EC}_{\text{bb}}$ ,  $\text{EC}_{\text{fossil}}$ ,  $\text{OC}_{\text{nf}}$  and  $\text{OC}_{\text{fossil}}$ , the derived mean represents the best estimate, and the SD represents the combined uncertainties (Table S3). For  $\text{POC}_{\text{bb}}$ ,  $\text{OC}_{\text{o,nf}}$ ,  $\text{POC}_{\text{fossil}}$ ,  $\text{SOC}_{\text{fossil}}$ ,  $\text{SOC}$  and  $f_{\text{fossil}}(\text{SOC})$ , the median value is considered as the best estimate and the interquartile ranges (25th–75th percentile) represent the combined uncertainties, because the PDFs of  $\text{POC}_{\text{fossil}}$  and  $\text{SOC}_{\text{fossil}}$  are asymmetric (Fig. S5, Table S4). The median values for  $\text{POC}_{\text{bb}}$  and  $\text{OC}_{\text{o,nf}}$  are very close to their mean values due to their symmetric PDFs (Fig. S5).

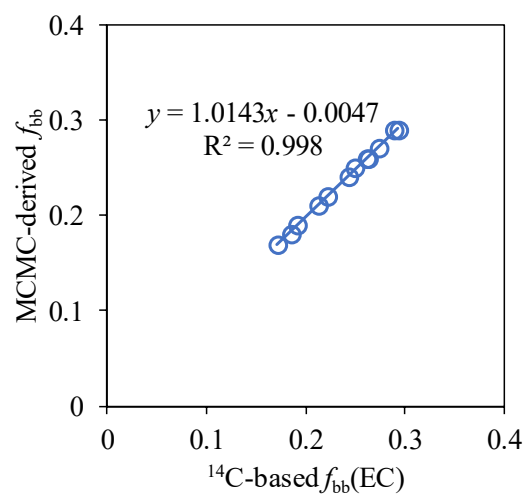


**Figure S1.** PM<sub>2.5</sub> mass concentrations (μg m<sup>-3</sup>) in Xi'an and Beijing during the measurement periods. The dashed areas indicate the selected samples for <sup>14</sup>C analysis. **(a)** 12h averaged PM<sub>2.5</sub> concentrations in Xi'an, the open symbols represent daytime samples, the filled square symbols represent nighttime samples. **(b)** 24h averaged PM<sub>2.5</sub> concentrations in Beijing. Samples selected for <sup>14</sup>C analysis are highlighted in red (indicating haze periods) and green (clean periods). For details, see Table S1.

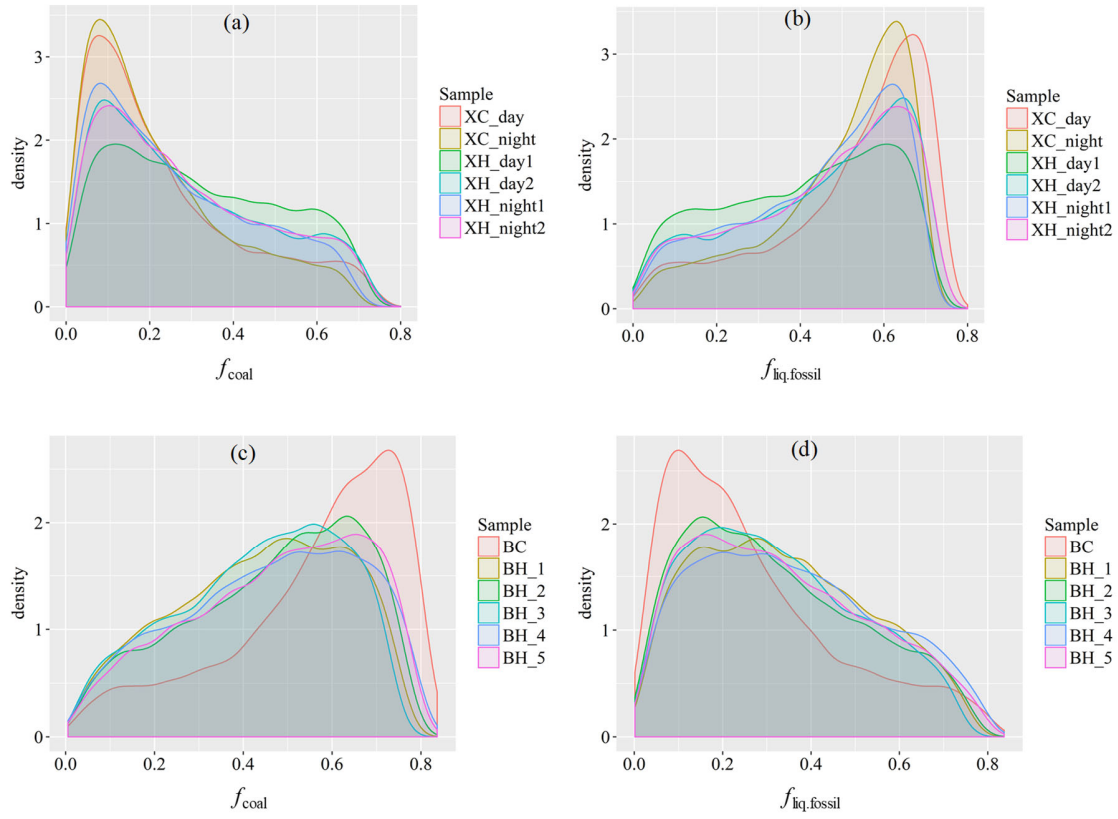


**Figure S2.** Temperature protocol of OC and EC isolation for  $^{14}\text{C}$  measurements was implemented on aerosol combustion system (Dusek et al., 2014) in pure  $\text{O}_2$ . OC is extracted by combusting the filter samples at 375 °C for 10 min. To isolate EC, OC is completely removed by 3 steps in order: water-extraction of filter samples (i.e., removal of water-soluble OC), combustion at 375 °C for 10 min (removal of water-insoluble OC; WIOC) in  $\text{O}_2$ , and combustion at 450 °C for 3 min in  $\text{O}_2$  (removal of the most refractory OC). Then, EC is isolated by heating the remaining carbon at 650 °C for 5 min in  $\text{O}_2$ . Details can be found in Zenker et al. (2017) and Ni et al. (2018).

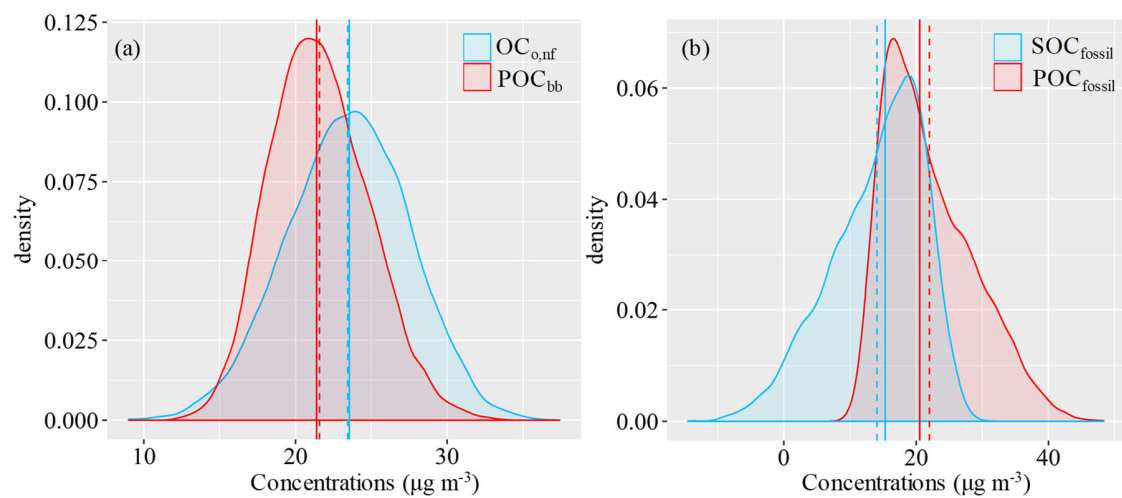




**Figure S3.** Comparison between the MCMC-derived fraction of biomass burning EC ( $f_{bb}$  derived from MCMC; median) and that obtained from radiocarbon data ( $^{14}\text{C}$ -based  $f_{bb}(\text{EC})$ ; mean).



**Figure S4.** Probability density functions (PDFs) of the relative source contribution of coal combustion ( $f_{\text{coal}}$ ) in Xi'an (a) and Beijing (c). PDFs of the relative source contribution of liquid fossil fuel combustion ( $f_{\text{liq.fossil}}$ ) in Xi'an (b) and Beijing (d).  $f_{\text{coal}}$  and  $f_{\text{liq.fossil}}$  are constrained by combining radiocarbon and  $\delta^{13}\text{C}$  measurements of EC, calculated using the Bayesian Markov chain Monte Carlo approach. For details, see Sect. 2.4.



**Figure S5. (a)** Example probability density functions (PDFs) of concentrations of  $\text{OC}_{\text{o,nf}}$  (light blue) and  $\text{POC}_{\text{bb}}$  (red) for sample XH\_day2. **(b)** PDFs of concentrations of  $\text{SOC}_{\text{fossil}}$  (light blue) and  $\text{POC}_{\text{fossil}}$  (red) for the same sample. Their concentrations are estimated from  $^{14}\text{C}$ -apportioned OC and EC using the EC tracer method (Sect. 2.4). The mean and median are indicated by the dashed and solid vertical lines, respectively.

**Table S1.** Details of sampling information and selected samples for radiocarbon measurements.

City	Descriptions	Selected samples for $^{14}\text{C}$ analysis			RH* (%)	Temperature (°C)	Wind speed (m s <sup>-1</sup> )
		Sample name	Note	Sampling Date/Time			
Xi'an	PM <sub>2.5</sub> samples were collected on the rooftop (~10 m) of a two-floor building located at the Institute of Earth Environment, Chinese Academy of Sciences (34.2° N, 108.9° E). This site is a typical urban background site surrounded by residential and education areas.	XH_day1	Haze/daytime	2017.1.1/Daytime 2017.1.2/Daytime	69 (56~91)	4.2 (-1.5~8.3)	1.0 (0.0~2.6)
		XH_night1	Haze/nighttime	2017.1.1/Nighttime 2017.1.2/Nighttime	88 (68~96)	0.3 (-2.6~3.5)	1.3 (0.2~2.5)
		XH_day2	Haze/daytime	2017.1.3/Daytime 2017.1.4/Daytime	78 (64~96)	4.4 (-2.6~7.2)	1.0 (0.0~2.3)
		XH_night2	Haze/nighttime	2017.1.3/Nighttime 2017.1.4/Nighttime	93 (84~97)	0.8 (-0.6~4.1)	0.6 (0.0~1.2)
		XC_day	Clean/daytime	2017.1.7/Daytime 2017.1.8/Daytime	60 (40~93)	4.4 (-4.2~8.0)	1.6 (0.4~2.6)
		XC_night	Clean/nighttime	2017.1.7/Nighttime 2017.1.8/Nighttime	85 (75~93)	-0.9 (-3.6~1.0)	1.0 (0.0~1.8)
Beijing	24 h integrated PM <sub>2.5</sub> samples were collected on the roof of a five-story building (~20 m) at the National Centre for Nanoscience (39.99 ° N, 116.32 ° E). The sampling site is close to the fourth ring of Beijing, and surrounded by residential, commercial and traffic areas.	BH_1	Haze	2016.12.11	59 (47~72)	0.7 (-1.7~2.1)	1.1 (0.4~1.7)
		BH_2	Haze	2016.12.17	66 (38~91)	-0.7 (-5.7~5.8)	0.8 (0.0~1.9)
		BH_3	Haze	2016.12.20	78 (55~88)	0.7 (-1.2~3.7)	1.1 (0.3~2.2)
		BH_4	Haze	2017.1.1 2017.1.2	59 (25~89)	0.0 (-5.7~8.2)	1.4 (0.3~3.1)
		BH_5	Haze	2017.1.4	73 (34~90)	1.1 (-3.0~7.3)	1.1 (0.0~1.7)
		BC	Clean	2016.12.26 2017.1.8	33 (22~48)	0.5 (-4.9~8.0)	1.8 (0.3~3.3)

\*The meteorological data (mean; minimum-maximum) is obtained from the Meteorological Institute of Shaanxi Province, Xi'an, China.

**Table S2.** Fraction modern ( $F^{14}C$ ) of elemental carbon (EC), organic carbon (OC) ( $F^{14}C_{(EC)}$  and  $F^{14}C_{(OC)}$ , respectively), and stable carbon isotopic compositions ( $\delta^{13}C$ , ‰) of EC ( $\delta^{13}C_{EC}$ ).

Sample name	Note	Sampling Date/Time	$F^{14}C_{(EC)}^a$	$F^{14}C_{(OC)}^a$	$\delta^{13}C_{EC}^a$
XH_day1	Haze/daytime	2017.1.1/Daytime 2017.1.2/Daytime	$0.301 \pm 0.007$	$0.594 \pm 0.004$	$-24.38 \pm 0.02$
XH_night1	Haze/nighttime	2017.1.1/Nighttime 2017.1.2/Nighttime	$0.321 \pm 0.014$	$0.608 \pm 0.005$	$-25.02 \pm 0.01$
XH_day2	Haze/daytime	2017.1.3/Daytime 2017.1.4/Daytime	$0.287 \pm 0.009$	$0.605 \pm 0.004$	$-24.92 \pm 0.04$
XH_night2	Haze/nighttime	2017.1.3/Nighttime 2017.1.4/Nighttime	$0.288 \pm 0.010$	$0.603 \pm 0.004$	$-24.87 \pm 0.03$
XC_day	Clean/daytime	2017.1.7/Daytime 2017.1.8/Daytime	$0.273 \pm 0.005$	$0.550 \pm 0.004$	$-25.53 \pm 0.02$
XC_night	Clean/nighttime	2017.1.7/Nighttime 2017.1.8/Nighttime	$0.317 \pm 0.004$	$0.590 \pm 0.004$	$-25.63 \pm 0.03$
BH_1	Haze	2016.12.11	$0.243 \pm 0.005$	$0.383 \pm 0.005$	$-24.43 \pm 0.03$
BH_2	Haze	2016.12.17	$0.233 \pm 0.005$	$0.368 \pm 0.005$	$-24.09 \pm 0.04$
BH_3	Haze	2016.12.20	$0.266 \pm 0.005$	$0.390 \pm 0.004$	$-24.38 \pm 0.02$
BH_4	Haze	2017.1.1 2017.1.2	$0.203 \pm 0.005$	$0.333 \pm 0.004$	$-24.29 \pm 0.02$
BH_5	Haze	2017.1.4	$0.211 \pm 0.003$	$0.334 \pm 0.003$	$-24.15 \pm 0.01$
BC	Clean	2016.12.26 2017.1.8	$0.188 \pm 0.006$	$0.403 \pm 0.005$	$-23.41 \pm 0.01$

<sup>a</sup> Values are given in average  $\pm$  measurement uncertainty.

**Table S3.** Fraction of non-fossil carbon in EC and OC ( $f_{bb}(EC)$ ,  $f_{nf}(OC)$ ), fraction of fossil carbon in EC and OC ( $f_{fossil}(EC)$ ,  $f_{fossil}(OC)$ ), concentrations ( $\mu\text{g m}^{-3}$ ) of EC and OC from non-fossil sources ( $EC_{bb}$  and  $OC_{nf}$ ) and fossil sources ( $EC_{fossil}$  and  $OC_{fossil}$ ) during haze and clean periods in Xi'an and Beijing during December 2016 and January 2017. Details of samples are shown in Table S1.

Sample name	$f_{bb}(EC)$	$f_{fossil}(EC)$	$f_{nf}(OC)$	$f_{fossil}(OC)$	$EC_{bb}$	$EC_{fossil}$	$OC_{nf}$	$OC_{fossil}$
XH_day1	$0.274 \pm 0.008$	$0.726 \pm 0.008$	$0.545 \pm 0.011$	$0.455 \pm 0.011$	$3.48 \pm 0.36$	$9.24 \pm 0.91$	$28.49 \pm 1.53$	$23.74 \pm 1.31$
XH_night1	$0.292 \pm 0.014$	$0.708 \pm 0.014$	$0.558 \pm 0.011$	$0.442 \pm 0.011$	$4.23 \pm 0.47$	$10.24 \pm 1.05$	$31.23 \pm 1.70$	$24.77 \pm 1.40$
XH_day2	$0.261 \pm 0.010$	$0.739 \pm 0.010$	$0.555 \pm 0.011$	$0.445 \pm 0.011$	$5.37 \pm 0.58$	$15.22 \pm 1.55$	$45.04 \pm 2.41$	$36.08 \pm 2.02$
XH_night2	$0.262 \pm 0.010$	$0.738 \pm 0.010$	$0.553 \pm 0.011$	$0.447 \pm 0.011$	$6.37 \pm 0.69$	$17.96 \pm 1.82$	$48.97 \pm 2.62$	$39.55 \pm 2.22$
XC_day	$0.248 \pm 0.006$	$0.752 \pm 0.006$	$0.505 \pm 0.010$	$0.495 \pm 0.010$	$2.17 \pm 0.23$	$6.56 \pm 0.66$	$12.61 \pm 0.68$	$12.37 \pm 0.66$
XC_night	$0.288 \pm 0.006$	$0.712 \pm 0.006$	$0.542 \pm 0.011$	$0.458 \pm 0.011$	$1.76 \pm 0.18$	$4.33 \pm 0.44$	$9.31 \pm 0.50$	$7.88 \pm 0.44$
BH_1	$0.221 \pm 0.006$	$0.779 \pm 0.006$	$0.352 \pm 0.008$	$0.648 \pm 0.008$	$3.04 \pm 0.32$	$10.71 \pm 1.08$	$15.02 \pm 0.83$	$27.66 \pm 1.44$
BH_2	$0.212 \pm 0.006$	$0.788 \pm 0.006$	$0.337 \pm 0.008$	$0.663 \pm 0.008$	$2.80 \pm 0.29$	$10.44 \pm 1.04$	$13.13 \pm 0.71$	$25.81 \pm 1.31$
BH_3	$0.242 \pm 0.006$	$0.758 \pm 0.006$	$0.358 \pm 0.008$	$0.642 \pm 0.008$	$3.97 \pm 0.41$	$12.42 \pm 1.25$	$18.62 \pm 1.01$	$33.43 \pm 1.72$
BH_4	$0.185 \pm 0.006$	$0.815 \pm 0.006$	$0.306 \pm 0.007$	$0.694 \pm 0.007$	$2.24 \pm 0.23$	$9.86 \pm 0.97$	$11.35 \pm 0.62$	$25.77 \pm 1.29$
BH_5	$0.191 \pm 0.004$	$0.809 \pm 0.004$	$0.306 \pm 0.007$	$0.694 \pm 0.007$	$2.31 \pm 0.24$	$9.74 \pm 0.97$	$12.23 \pm 0.67$	$27.68 \pm 1.42$
BC	$0.171 \pm 0.006$	$0.829 \pm 0.006$	$0.370 \pm 0.008$	$0.630 \pm 0.008$	$0.27 \pm 0.03$	$1.33 \pm 0.13$	$2.54 \pm 0.14$	$4.34 \pm 0.23$

**Table S4.** Concentrations ( $\mu\text{g m}^{-3}$ ) of primary OC from biomass burning ( $\text{POC}_{\text{bb}}$ ), primary OC from fossil sources ( $\text{POC}_{\text{fossil}}$ ), OC from non-fossil sources excluding primary biomass burning ( $\text{OC}_{\text{o,nf}}$ ), secondary OC from fossil sources ( $\text{SOC}_{\text{fossil}}$ ) and total SOC (i.e., approximately the sum of  $\text{OC}_{\text{o,nf}}$  and  $\text{SOC}_{\text{fossil}}$ ) during haze and clean periods in Xi'an and Beijing during December 2016 and January 2017 (median and interquartile range). Details of samples are shown in Table S1.

Sample name	$\text{POC}_{\text{bb}}$	$\text{POC}_{\text{fossil}}$	$\text{OC}_{\text{o,nf}}$	$\text{SOC}_{\text{fossil}}$	SOC	$f_{\text{fossil}}(\text{SOC})$
XH_day1	13.82 (12.50–15.26)	13.38 (10.61–16.88)	14.58 (12.83–16.25)	10.34 (6.71–13.16)	24.56 (20.53–28.25)	0.41 (0.32–0.47)
XH_night1	16.80 (15.14–18.58)	13.80 (11.24–17.61)	14.40 (12.27–16.41)	10.82 (7.04–13.61)	24.81 (20.35–28.80)	0.43 (0.33–0.49)
XH_day2	21.43 (19.28–23.68)	20.50 (16.60–26.58)	23.54 (20.75–26.20)	15.41 (9.35–19.52)	38.29 (31.42–44.04)	0.40 (0.29–0.45)
XH_night2	25.40 (22.88–28.07)	24.32 (19.66–30.96)	23.51 (20.34–26.62)	15.06 (8.37–19.94)	37.85 (30.55–44.63)	0.39 (0.27–0.46)
XC_day	8.64 (7.77–9.54)	8.21 (6.82–10.63)	3.98 (2.95–4.94)	4.07 (1.73–5.59)	7.76 (5.18–9.89)	0.52 (0.36–0.60)
XC_night	6.99 (6.30–7.72)	5.37 4.46–6.85	2.32 1.51–3.07	2.47 1.00–3.43	4.57 2.83–6.07	0.53 0.38–0.64
BH_1	12.15 (10.90–13.44)	18.32 (14.82–21.65)	2.87 (1.47–4.27)	9.44 (5.86–12.99)	12.33 98.13–16.41)	0.76 (0.67–0.86)
BH_2	11.17 (10.04–12.34)	18.36 (14.89–21.44)	1.98 (0.65–3.17)	7.56 (4.26–11.04)	9.59 (5.67–13.45)	0.80 (0.68–0.91)
BH_3	15.82 (14.21–17.56)	21.21 (17.18–25.02)	2.77 (0.95–4.50)	12.24 (8.16–16.36)	15.01 (10.19–19.75)	0.81 (0.72–0.92)
BH_4	8.88 (8.00–9.84)	16.95 (13.83–19.94)	2.45 (1.41–3.44)	8.79 (5.70–12.06)	11.30 (7.71–14.83)	0.78 (0.70–0.86)
BH_5	9.17 (8.25–10.17)	16.94 (13.73–19.86)	3.00 (1.96–4.07)	10.79 (7.65–14.08)	13.82 (10.26–17.53)	0.78 (0.71–0.85)
BC	1.09 (0.98–1.21)	2.55 (2.15–2.90)	1.44 (1.30–1.59)	1.80 (1.42–2.22)	3.26 (2.79–3.76)	0.55 (0.50–0.60)

**Table S5.** Fractional contribution of different incomplete combustion sources to EC in different [seasons-cities](#) ( $f_{bb}$ ,  $f_{liq.fossil}$  and  $f_{coal}$ ; median and interquartile range ), and EC concentrations ( $\mu\text{g m}^{-3}$ ) from biomass burning ( $EC_{bb}$ ), coal combustion ( $EC_{coal}$ ) and liquid fossil fuel combustion ( $EC_{liq.fossil}$ ) (median and interquartile range). Details of samples are shown in Table S1.

City	Sample Name	Note	$f_{bb}$	$f_{liq.fossil}$	$f_{coal}$	$EC_{bb}$	$EC_{liq.fossil}$	$EC_{coal}$
Xi'an	XH_day1	Haze1/daytime	0.27 (0.27–0.28)	0.44 (0.26–0.58)	0.29 (0.14–0.47)	3.48 (3.24–3.73)	5.49 (3.10–7.30)	3.66 (1.82–5.98)
	XH_night1	Haze1/nighttime	0.29 (0.29–0.30)	0.47 (0.30–0.59)	0.23 (0.12–0.41)	4.23 (3.93–4.54)	6.75 (4.28–8.45)	3.29 (1.69–5.87)
	XH_day2	Haze2/daytime	0.29 (0.29–0.30)	0.47 (0.30–0.59)	0.23 (0.12–0.41)	5.38 (4.99–5.78)	10.04 (6.15–12.60)	4.94 (2.44–8.94)
	XH_night2	Haze2/nighttime	0.26 (0.26–0.27)	0.49 (0.30–0.62)	0.25 (0.12–0.44)	6.39 (5.94–6.85)	11.85 (7.40–14.93)	5.88 (2.91–10.35)
	XC_day	Clean/daytime	0.25 (0.24–0.26)	0.57 (0.40–0.66)	0.18 (0.09–0.35)	2.17 (2.01–2.32)	4.73 (3.23–5.68)	1.69 (0.83–3.24)
	XC_night	Clean/nighttime	0.29 (0.28–0.30)	0.54 (0.38–0.62)	0.17 (0.09–0.33)	1.76 (1.63–1.89)	3.17 (2.21–3.76)	1.07 (0.53–2.05)
Beijing	BH_1	Haze1	0.22 (0.21–0.23)	0.32 (0.18–0.49)	0.46 (0.29–0.60)	3.05 (2.83–3.28)	4.38 (2.49–6.68)	6.18 (3.90–8.18)
	BH_2	Haze2	0.21 (0.21–0.22)	0.29 (0.16–0.46)	0.50 (0.33–0.62)	2.80 (2.59–3.02)	3.89 (2.15–6.13)	6.42 (4.20–8.15)
	BH_3	Haze3	0.24 (0.23–0.25)	0.31 (0.18–0.47)	0.45 (0.29–0.58)	3.98 (3.69–4.28)	5.14 (2.99–7.79)	7.17 (4.52–9.38)
	BH_4	Haze4	0.19 (0.18–0.19)	0.33 (0.18–0.50)	0.49 (0.31–0.63)	2.24 (2.07–2.40)	3.90 (2.21–6.00)	5.80 (3.75–7.53)
	BH_5	Haze5	0.19 (0.19–0.20)	0.31 (0.17–0.48)	0.50 (0.33–0.63)	2.31 (2.13–2.48)	3.75 (2.09–5.87)	5.86 (3.73–7.55)
	BC	Clean	0.17 (0.17–0.18)	0.22 (0.12–0.37)	0.61 (0.46–0.71)	0.27 (0.25–0.30)	0.35 (0.18–0.59)	0.96 (0.72–1.14)



**Table S6.** Fractional contribution of different incomplete combustion sources to EC (median and interquartile range).

	Biomass burning	Coal combustion	Liquid fossil fuel combustion (i.e., vehicle emissions)
Xi'an haze day	0.27 (0.26–0.27)	0.26 (0.13–0.44)	0.47 (0.29–0.60)
Xi'an haze night	0.28 (0.27–0.28)	0.23 (0.12–0.41)	0.49 (0.31–0.61)
Xi'an clean day	0.25 (0.24–0.26)	0.18 (0.09–0.33)	0.57 (0.42–0.66)
Xi'an clean night	0.29 (0.28–0.30)	0.18 (0.09–0.34)	0.53 (0.37–0.62)
Beijing haze	0.21 (0.20–0.22)	0.48 (0.31–0.61)	0.31 (0.18–0.48)
Beijing clean	0.17 (0.17–0.18)	0.61 (0.45–0.71)	0.22 (0.12–0.38)

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