We thank the reviewers for the helpful comments and providing us the opportunity to strengthen our research. We try to address all of them carefully. Below are point-to-point responses.

#### **Anonymous reviewer #1:**

The measurement report by Ni et al. is an excellent manifestation of the measurement report paper. Despite not presenting strikingly new results, comparative studies involving lesser studied cities or regions are encouraged as long as the data analysis is done robustly and extensively. I particularly commend Bayesian MCMC simulation in addition to <sup>14</sup>C analysis which enhances the results as well as excellent and informative Figures. The paper is very well written and can be published after addressing few minor comments.

1) Line 22. The term "clean" can hardly be attributed to mega-city environment. Perhaps "moderately polluted" or at least "relatively clean". Haze is typically related to poor visibility, so perhaps the use of "clear weather", which can be quite polluted too, can be more informative. The term "clean" is used throughout the text and I encourage finding a proper substitute.

**Response:** We agree with the reviewer that the "clean" periods we defined in this manuscript cannot fully represent the pollution levels, especially for "clean" periods in Xi'an. In Xi'an, as stated in Sect. 3.1 (page 6, line 172–173), even during clean periods we defined here, daily PM<sub>2.5</sub> concentrations (<100  $\mu$ g m<sup>-3</sup>) were higher than the Chinese pollution standard of 75  $\mu$ g m<sup>-3</sup>. However, in this study, we intent to use "clean" and "haze" to represent low and high pollution events, respectively, classified by PM<sub>2.5</sub> concentrations. With this study scope, we would prefer not to use the term "clear weather", because it is a weather condition, which cannot be identified by PM<sub>2.5</sub> concentrations.

In this study, we define clean days as daily average  $PM_{2.5}$  concentrations < median concentrations in the winter heating season (15 November 2016 to 15 March 2017). If possible, we think it probably acceptable to keep the term "clean" to facilitate the writing, and to avoid the inconsistence between the discussion paper and the final one (e.g., "clean" are included in nearly all figures and tables as legend, text annotation etc.). In the revised manuscript, we explain quantitively the definition of "clean" in this study (changes are underlined):

(i) Abstract lines 16–19 (page 1)

"Here we combined the analysis of radiocarbon and the stable isotope <sup>13</sup>C to investigate the sources and formation of carbonaceous aerosols collected in two Chinese megacities (Beijing and Xi'an) during severe haze events of "red alarm" level from December 2016 to January 2017. The haze periods with daily  $PM_{2.5}$  concentrations as high as ~400 µg m<sup>-3</sup> were compared to subsequent clean periods (i.e.,  $PM_{2.5}$  < median concentrations during the winter 2016/2017), with  $PM_{2.5}$  concentrations below 100 µg m<sup>-3</sup> in Xi'an and below 20 µg m<sup>-3</sup> in Beijing."

(ii) Manuscript lines 101–103 (page 4)

Six samples from haze and clean days were selected per sampling site for carbon isotope analysis (Tables S1 and S2). We define clean days at each site as  $PM_{2.5}$  < median concentration in the winter heating season from 15 November 2016 to 15 March 2017.

(iii) Figure caption (page 18)

**Figure 1.** Mass concentrations of OC and EC from fossil and non-fossil sources (OC<sub>fossil</sub>, OC<sub>nf</sub>, EC<sub>fossil</sub> and EC<sub>bb</sub>) as well as fraction of non-fossil carbon in OC and EC ( $f_{nf}$ (OC) and  $f_{bb}$ (EC), respectively) for daytime and nighttime PM<sub>2.5</sub> samples in Xi'an, and 24h-integrated PM<sub>2.5</sub> samples in Beijing <u>during haze and clean periods during the measurement periods (2</u> December 2016 to 10 January 2017). For each city "haze" and "clean" are used to represent high and low pollution events, and clean days at each site are defined as days with PM<sub>2.5</sub> < median concentration in the winter heating season from 15 November 2016 to 15 March 2017. Uncertainties of <sup>14</sup>C-apportioned  $f_{nf}$ (OC) and  $f_{bb}$ (EC) are indicated but are too small to be visible. The data are shown in Table S3.

2) Line 33. Field measurements.

Response: Thank you for spotting out the typo. Corrected (page 2, line 36).

"<u>Field</u> measurements show that carbonaceous aerosol contributes a significant fraction of PM<sub>2.5</sub> loading during severe haze events in China"

**3**) Line 35. I think it should be stressed that better understanding of sources outside capital Beijing is needed for comparison as well as for more comprehensive understanding. Beijing was fairly well studied already.

**Response:** We improved the introduction according to the comments and suggestions by the reviewer. The following underlined sentences are added in the introduction (page 2, line 39–41):

"Severe haze pollution with high  $PM_{2.5}$  (i.e., particulate matter with aerodynamic diameter  $\leq 2.5 \ \mu$ m) concentrations and reduced visibility occurs frequently during winter in China (An et al., 2019). Field measurements show that carbonaceous aerosol contributes a significant fraction of  $PM_{2.5}$  loading during severe haze events in China (Huang et al., 2014; Elser et al., 2016; Liu et al., 2016). Therefore, a better understanding of the sources and atmospheric processes of carbonaceous aerosols is needed for mitigating haze pollution. Many previous studies focus solely on Beijing, the capital of China. However, studies on other megacities are also needed for comparison as well as for a more comprehensive understanding of haze pollution in China."

4) Line 85. Please indicate similarity quantitatively as there is a proportionality issue commented later.Response: The revised text shows:

"Each composite sample consists of 2 12h (for Xi'an) or 24 h (for Beijing) filter pieces with similar  $PM_{2.5}$  loading that agree within 20 % (Fig. S1)." (page 4, line 105–106)

**5**) Line 149. Concentrations neither <100 nor <20 can be considered clean, especially that the two numbers differ by five times. If haze concentrations in Xian were defined >250, that is only 2-3 times different to clean, so certainly qualifies for moderate pollution. Furthermore, if Chinese national pollution standard is at 75ug/m3, concentration <100 can in no way qualify for clean.

**Response:** Thank you for pointing this out. In this study, the haze periods with daily  $PM_{2.5}$  concentrations as high as ~400 µg m<sup>-3</sup> were compared to subsequent clean periods, with  $PM_{2.5}$  concentrations below 100 µg m<sup>-3</sup> in Xi'an and below 20 µg m<sup>-3</sup> in Beijing. We define clean days as daily average  $PM_{2.5}$  concentrations < median concentrations in the winter heating season (15 November 2016 to 15 March 2017). This has been clarified in the revised manuscript, as addressed in the response to question 1.

We agree with the reviewer that in Xi'an  $PM_{2.5}$  concentrations< 100 µg m<sup>-3</sup> quantifies "moderate pollution", as state in the Sect. 3.1: <u>"In Xi'an, even during clean periods we defined here, daily  $PM_{2.5}$  concentrations were higher than the Chinese pollution standard of 75 µg m<sup>-3</sup>, reflecting severe air quality problems." (page 6, line 172–173). However,  $PM_{2.5}$  concentration < 20 µg m<sup>-3</sup> in Beijing can be considered as clean, because it is smaller than the 24 h mean of 25 µg m<sup>-3</sup> suggested by the Air Quality Guidelines of World Health Organization (WHO, 2006).</u>

In the revised manuscript, to avoid any confusing, the term "clean" is defined quantitively as daily average  $PM_{2.5}$  concentrations < median concentrations in the winter heating season (15 November 2016 to 15 March 2017), and is explained and clarified several times, for example, in the abstract and in the main text when "clean" is used (see changes in response to question 1).

6) Line 177. will be much higher.

Response: Done (page 7, line 199).

"So even though biomass burning contributes a small portion of EC, its contribution to primary OC will <u>be</u> much higher."

7) Line 250. ... emissions are very low in winter.

Response: Done (page 9, line 279).

"The  $OC_{o,nf}$  is not likely attributed to biogenic OC, because the biogenic emissions are very low <u>in winter.</u>"

**8**) Line 277. Why would large secondary formation from fossil sources be particularly favoured in Beijing only? With no plausible explanation it should be dismissed. Stronger contribution from coal combustion (both primary and secondary) sounds convincing given Beijing geographical location.

**Response:** In this study, large secondary formation from fossil sources in Beijing is justified by larger  $f_{\text{fossil}}(\text{SOC})$  in Beijing than in Xi'an (Fig. 6), where  $f_{\text{fossil}}(\text{SOC})$  is the fraction of fossil SOC in total SOC of both fossil and non-fossil origins. However, the explanations for large secondary formation from fossil sources in Beijing are not clear yet, as the reviewer pointed out. To clarify, the following underlined sentences are added in the revised manuscript:

"Compared to Xi'an, Beijing had significantly smaller differences between  $f_{bb}(EC)$  and  $f_{nf}(OC)$ (Fig. 1), which was also observed in previous studies during the haze event in January 2013 (Zhang et al., 2015). <u>Comparing Eq. (14) with Eq. (16)</u>, this suggests either strong contribution from coal combustion in Beijing or large secondary formation from fossil sources, or both. The stronger contribution of coal combustion to OC in Beijing than in Xi'an was a direct consequence of a larger proportion of coal combustion in EC in Beijing, as demonstrated by the Bayesian MCMC results of EC (Sect. 3.2). The latter was further validated by the variation of SOC. The  $f_{fossil}(SOC)$  in Beijing was higher than that in Xi'an, despite the variations between haze and clean periods (Sect. 3.3). <u>By combining <sup>14</sup>C</u> measurements with other state-of-art analytical techniques (e.g., aerosol mass spectrometry), Huang et al. (2014) also found that fossil OC was mostly secondary in nature in Beijing, and non-fossil SOC formation was dominant in Xi'an during a wintertime haze episode (i.e., Beijing had a larger  $f_{fossil}(SOC)$  than Xi'an). However, atmospheric mechanisms responsible for the enhancement in fossil-derived SOA formation in Beijing remain unclear." (page 11, line 316–319)

9) Line 287. ... by significantly larger...

**Response:** We have corrected the text as suggested:

"This is validated by <u>significantly larger</u>  $f_{\text{fossil}}(\text{SOC})$  during haze periods (76%–81%) than during clean periods (~55%; Sect. 3.3)." (page 11, line 326).

10) Line 294. ... wide range of EC concentrations.

Response: Corrected.

"In Xi'an,  $f_{bb}(EC)$  was fairly constant during haze and clean periods, despite the wide <u>range</u> of EC concentrations." (page 12, line 336)

**11**) Line 301. same as above comment (i.e., question 8)

**Response:** Due to the unclear mechanisms responsible for the enhancement in fossil SOC in Beijing (as addressed in response to question 8), we would like to avoid to highlight this point in the conclusion, thus delete the following underlined sentence in the conclusion of the revised manuscript:

"The differences between non-fossil fraction in OC and EC were smaller in Beijing and larger in Xi'an. <u>This suggests strong contribution from coal combustion or large secondary</u> <u>formation from fossil sources, or both in Beijing.</u>" (page 12, line 343–344)

## **References:**

WHO: Air Quality Guidelines: Global Update 2005: Particulate Matter, Ozone, Nitrogen Dioxide and Sulfur Dioxide, World Health Organization, 2006.

#### Anonymous reviewer #2:

The article by Ni et al entitled "Measurement report: Dual-carbon isotopic characterization of carbonaceous aerosol in Beijing and Xi'an: distinctions in primary versus secondary sources" is presenting application of natural carbon isotopes to characterize carbonaceous aerosol composition in Beijing and Xi'an (two Chinese mega-cities). The title of the paper does not seem to fully reflect the subject of the article. In fact, the spectrum of issues discussed in the paper is much wider than "distinctions in primary versus secondary sources". While we are observing an increased interest in the topics related to the study of air quality and its impact on the health of citizens, the development of various methods dedicated for identification the pollution sources is a very important topic that fits into the scope of the journal. Carbonaceous aerosols constitute an important fraction of air pollution observed in mega-cities. Application of carbon isotopes allows to identify the share and its temporal and spatial dynamics of different emission sources. In this context the paper address relevant scientific problem and demonstrates a possible solution based on comprehensive use of isotopic tracers applied for both elemental carbon (EC) and organic carbon (OC) fractions of carbonaceous pollutants. The presented methodology was applied for the short-term measurement campaigns performed in two Chinese mega-cities having different emission sources structure. Authors demonstrated usefulness of this methodology for identifying share of the pollutants having different origin and presented an interesting data increasing the understanding of the differences between at first glance similar urban environments. The paper is well structured, the description of methodology is clear and, in my opinion, complete. Abstract contains a clear message of the paper. The results are correctly presented and, in most parts, well discussed. The quality and number of figures is correct. Presented results are discussed in relation to other studies appropriate referenced in the text. I recommend to publish the paper after a minor revision.

**Response:** Following the reviewer's concern about the title, we change the title to better reflect the content of the manuscript and to be more specific (changes are underlined):

## "Dual-carbon isotopic characterization of carbonaceous aerosol <u>reveal different</u> primary and secondary sources in Beijing and Xi'an during severe haze events"

We would like to avoid including "day-night difference" in the title, because it is only studied in Xi'an. Furthermore, the day-night difference in sources of carbonaceous aerosol was slight in this study, especially during haze periods in Xi'an.

### **Specific comments**

1) The authors discuss in the paper a wide spectrum of issues including different contribution of coal, liquid fossil fuel and biomass combustion to elemental and organic fraction of carbonaceous aerosols, temporal (day-time vs. night time and haze events vs. clean periods) variability as well as spatial

(location specific differences between Beijing and Xi'an cities) variability. In addition, issues related to primary and secondary organic aerosols are discussed. In such wide range of discussed aspects, it is difficult to keep the description clear. Maybe a short summary following each part of the discussion would help reader to keep on track of the analysis.

Response: The revised manuscript adds a short summary at the end of each result section.

(i) Sect. 3.1 Fossil and non-fossil contributions to EC and OC

"Overall, our <sup>14</sup>C data show that fossil sources contribute more strongly to EC and OC in Beijing than in Xi'an, which is consistent with previous observations. Both in Beijing and Xi'an the fossil vs. non-fossil contributions to EC and OC did not change drastically during haze and clean periods. In Xi'an, a slight day-night difference was observed during clean periods, but disappeared during haze periods, suggesting a large accumulation of particles." (page 7, line 215–218)

(ii) Sect. 3.2 Fossil EC apportioned by stable carbon isotopes: coal vs. liquid fossil fuel

"In summary, complementing <sup>14</sup>C with  $\delta^{13}$ C allows for quantitative constraints on EC sources: EC was dominated by liquid fossil fuel combustion (i.e., vehicle emissions) in Xi'an and by coal burning in Beijing, especially during clean periods. In Xi'an, no strong day-night differences in EC sources were observed during haze and clean periods." (page 9, line 248– 250)

(iii) Sect. 3.3 Primary and secondary OC

"Together these results reveal the differences in primary and secondary OC in two Chinese megacities. The contribution of SOC to total OC increased from clean to haze periods in Xi'an. In contrary, SOC/OC ratios increased from haze to clean periods in Beijing, mainly due to increased SOC from non-fossil sources. SOC was dominated by fossil sources in Beijing but by non-fossil sources in Xi'an, especially during haze periods. In Xi'an, the day-night difference was larger during clean periods with less SOC at night." (page 10, line 289–293)

(iv) Sect 3.4 Difference between the fractions of non-fossil carbon in OC and EC

"In conclusion, this section discusses the factors that governs the differences between  $f_{bb}(EC)$  and  $f_{nf}(OC)$ , and concludes that smaller differences suggest stronger contribution from coal combustion and/or larger secondary formation from fossil sources. This is further examined and validated by source apportionment results of EC and OC (Sect. 3.1–3.3)." (page 11, line 328–330)

**2**) The Methodology section is very detailed and well referenced but I found no information concerning possible mineral contamination (carbonates) of the collected samples is present. Were there any corrections to the mineral fraction contamination in the samples applied?

**Response:** Thank you for pointing this out. For fine particle  $PM_{2.5}$ , the interference of carbonate carbon with the OC and EC is negligible in most ambient (Chow and Watson, 2002; Querol et al., 2004), because carbonate is mostly in the coarse mode.

In this study, we did preliminary studies to check the possible interference of carbonate carbon.  $PM_{2.5}$  samples with the highest calcium concentrations (associate with carbonate in most cases; Chow and Watson, 2002) in Xi'an, China were used to check the presence of carbonate carbon. Carbonate carbon in an aerosol sample can be verified if there is significant difference in TC (i.e., total carbon) mass before and after acidification (i.e., expose the sample to HCl vapor; NIOSH, 1999). Negligible carbonate carbon was found, because TC mass before and after acidification agrees each other within the measurement uncertainty.  $PM_{2.5}$  samples in Beijing are less likely affected by carbonates, owing to lower dust concentrations (i.e., a major source of carbonate) than in Xi'an (Huang et al., 2014). In this study, we therefore did not correct the mineral contamination (carbonates), because the contamination is very small compared to the relatively larger OC and EC amounts for both mass determination and carbon isotopic analysis.

We add the above discussion in the Supplement S1. In the main text, we add (page 4, line 97–99):

"Acidification to remove potential interferences from carbonates is not necessary, because carbonate carbon in PM<sub>2.5</sub> samples is found to be negligible, compared to the relatively larger OC and EC amounts for both mass determination and carbon isotopic analysis (Supplement S1)."

**3**) More detailed description of study sites (location, topography, typical emissions) would help to understand the differences in presented results between two cities.

**Response:** In the revised manuscript, we add the site descriptions in the method section (page 3, line 76–88):

"Xi'an is the largest city in northwestern China, with over 8.8 million residents and 2.5 million vehicles in 2016 (Xi'an Municipal Bureau of Statistics and NBS Survey Office in Xi'an, 2017). Surrounded by Qinling Mountains to the south and the Loess Plateau to the north, days with low wind speed occur frequently in Xi'an, promoting the accumulation of air pollutants. Xi'an is now facing increased serious air quality issues due to the rapid increase of motor vehicles and energy consumption in the past two decades. Besides residential coal combustion, biomass burning is also a major emission source in Xi'an and its surrounding areas (i.e., Guanzhong Plain) for heating and cooking especially in winter (Zhang et al., 2014; Xu et al.,

2016). Beijing, the capital of China, is a megacity with over 21 million residents and 5.7 million vehicles in 2016 (Beijing Municipal Bureau of Statistics and NBS Survey Office in Beijing, 2017). Beijing is located in the Beijing-Tianjin-Hebei region, the most economically developed region in North China. However, the rapid economic growth and urbanization associated with heavy coal consumption and rapid increase usage of vehicles lead to the poor air quality in Beijing. Besides local emissions, regional transport of pollutants between neighboring cities also contributes to air pollution in Beijing (Zheng et al., 2015; An et al., 2019)."

The new citations are included in the revised reference list.

**4)** Did the authors consider the admixture of bio-fuels into liquid fossil component in the context of  $F^{14}C_{liq,fossil}$  parameter? Is it a case of Chinese liquid fuels market?

**Response:**  $F^{14}C_{liq.fossil}$  denotes  $F^{14}C$  of liquid fossil fuel, used in Eq. (13) for source apportionment of EC (page 6). In China,  $F^{14}C_{liq.fossil}$  has been approximated as zero in EC source apportionment studies (e.g., Andersson et al., 2015; Fang et al., 2018), because liquid fossil fuel combustion emits predominantly fossil carbon. However, in recent years a small fraction of bio-fuels has been added to liquid fossil fuel (e.g., gasoline and diesel) for on-road transportation. Nowadays, most of the biofuels are produced by USA (mainly corn ethanol) and Brazil (mainly sugarcane ethanol). The admixture of bio-fuels into liquid fossil fuel increases  $F^{14}C_{liq.fossil}$ , although it is still unclear to what extent it affects  $F^{14}C_{liq.fossil}$ .

Back to the winter 2016/2017 (the measurement period in this study), Beijing and Xi'an were not in the 11 provinces that adopted the fuel E10 gasoline (i.e., 10% biomass-based ethanol and 90% gasoline), the major bio-fuel for transportation in China (Hao et al., 2018). The development of bio-fuels in the transport sector in China has been restrained by food security concern and land availability (Zhang and Chen, 2015), because the bio-fuel in China is mainly derived from corn. EC from bio-fuels is a very minor fraction of EC in China (Zheng et al., 2019), hence it is not be considered in this study. This might change in the future if bio-fuels are used more extensively.

5) I don't see any clear reason for introducing equations 14 to 16.

**Response:** Equations 14 to 16 are introduced for illustrating what governs the differences between  $f_{bb}(EC)$  and  $f_{nf}(OC)$ . Comparing Eq. (14) with Eq. (16), we tried to explain that (a) the differences between  $f_{bb}(EC)$  and  $f_{nf}(OC)$  depend on the differences between  $\frac{(1+SOC_{fossil}/POC_{fossil})}{(1+OC_{o,nf}/POC_{bb})} \times \frac{r_{fossil}}{r_{bb}}$  and 1 (page 10, line 303), and (b) further elaborate on the statements that "With smaller  $r_{fossil}$  than  $r_{bb}$  similar  $f_{nf}(OC)$  and  $f_{bb}(EC)$  can result from larger secondary formation from fossil sources than non-fossil sources (i.e.,  $\frac{(1+SOC_{fossil}/POC_{fossil})}{(1+OC_{o,nf}/POC_{bb})} > 1$ .)" (page 11, line 305–306) and "Therefore, in a city where

biomass burning and coal combustion are the dominant pollution sources,  $f_{nf}(OC)$  and  $f_{bb}(EC)$  will be more similar than in a city where the main sources are biomass burning and vehicle emissions." (page 11, line 307–309). The underlined statements are not straightforward without introducing equations 14 to 16. Those equations are given, because they are not so common. Therefore, if the reviewer allows, we prefer to keep those equations.

### **References:**

Andersson, A., Deng, J., Du, K., Zheng, M., Yan, C., Sköld, M., and Gustafsson, Ö.: Regionally-varying combustion sources of the January 2013 severe haze events over eastern China, Environ. Sci. Technol., 49, 2038–2043, 2015.

Chow, J. C. and Watson, J. G.: PM<sub>2.5</sub> carbonate concentrations at regionally representative Interagency Monitoring of Protected Visual Environments sites, *J. Geophys. Res.*, 107( D21), 8344, doi:10.1029/2001JD000574, 2002.

Fang, W., Andersson, A., Zheng, M., Lee, M., Holmstrand, H., Kim, S.-W., Du, K., and Gustafsson, Ö.: Divergent evolution of carbonaceous aerosols during dispersal of East Asian haze, Sci. Rep.-UK, 7, 10422, <u>https://doi.org/10.1038/s41598-017-10766-4</u>, 2017.

Hao, H., Liu, Z., Zhao, F., Ren, J., Chang, S., Rong, K., and Du, J.: Biofuel for vehicle use in China: Current status, future potential and policy implications, Renew. Sust. Energ. Rev., 82, 645-653, <u>https://doi.org/10.1016/j.rser.2017.09.045</u>, 2018.

NIOSH: Method 5040 Issue 3 (Interim): Elemental carbon (diesel exhaust), in: NIOSH Manual of Analytical Methods; National Institute of Occupational Safety and Health: Cincinnati, OH, 1999.

Querol, X., Alastuey, A., Ruiz, C. R., Arti<sup>\*</sup>nano, B., Hansson, H. C., Harrison, R. M., Buringh, E., ten Brink, H. M., Lutz, M., Bruckmann, P., Straehl, P., and Schneider, J.: Speciation and origin of PM<sub>10</sub> and PM<sub>2.5</sub> in selected European cities, Atmos. Environ., 38, 6547–6555, 2004.

Zhang, H., and Chen, W.: The Role of Biofuels in China's Transport Sector in Carbon Mitigation Scenarios, Energy Proced., 75, 2700-2705, <u>https://doi.org/10.1016/j.egypro.2015.07.682</u>, 2015.

Zheng, X., Wu, X., He, L., Guo, X., and Wu, Y.: Black Carbon Emissions from Light-duty Passenger Vehicles Using Ethanol Blended Gasoline Fuels, Aerosol Air Qual. Res., 19, 1645-1654, 10.4209/aaqr.2019.02.0095, 2019.

## Measurement report: Dual-carbon isotopic characterization of carbonaceous aerosol in <u>Beijing and Xi'anreveal</u> : <u>distinctionsin</u> <u>different primary versus and secondary sources in Beijing and Xi'an</u> <u>during severe haze events</u>

5 Haiyan Ni<sup>1,2</sup>, Ru-Jin Huang<sup>1,3</sup>, Max M. Cosijn<sup>2</sup>, Lu Yang<sup>1</sup>, Jie Guo<sup>1</sup>, Junji Cao<sup>1</sup>, and Ulrike Dusek<sup>2</sup>

<sup>1</sup>State Key Laboratory of Loess and Quaternary Geology, Center for Excellence in Quaternary Science and Global Change, Key Laboratory of Aerosol Chemistry and Physics, Institute of Earth Environment, Chinese Academy of Sciences, Xi'an, 710061, China

<sup>2</sup>Centre for Isotope Research (CIO), Energy and Sustainability Research Institute Groningen (ESRIG), University of 10 Groningen, Groningen, 9747 AG, the Netherlands

<sup>3</sup>Institute of Global Environmental Change, Xi'an Jiaotong University, Xi'an, 710049, China

Correspondence: Ru-Jin Huang (rujin.huang@ieecas.cn)

Abstract. To mitigate haze pollution in China, a better understanding of the sources of carbonaceous aerosols is required due to the complexity in multiple emissions and atmospheric processes. Here we combined the analysis of radiocarbon and the

- 15 stable isotope <sup>13</sup>C to investigate the sources and formation of carbonaceous aerosols collected in two Chinese megacities (Beijing and Xi'an) during severe haze events of "red alarm" level from December 2016 to January 2017. <u>The haze periods</u> with daily PM<sub>2.5</sub> concentrations as high as ~400 µg m<sup>-3</sup> were compared to subsequent clean periods (i.e., PM<sub>2.5</sub> < median concentrations during the winter 2016/2017), with PM<sub>2.5</sub> concentrations below 100 µg m<sup>-3</sup> in Xi'an and below 20 µg m<sup>-3</sup> in Beijing. In Xi'an, liquid fossil fuel combustion was the dominant source of elemental carbon (EC; 44%–57%), followed by
- biomass burning (25%–29%) and coal combustion (17%–29%). In Beijing, coal combustion contributed 45%–61% of EC and biomass burning (17%–24%) and liquid fossil fuel combustion (22%–33%) contributed less. Non-fossil sources contributed 51%–56% of organic carbon (OC) in Xi'an and fossil sources contributed 63%–69% of OC in Beijing. Secondary OC (SOC) was largely contributed by non-fossil sources in Xi'an (56 ± 6%) and by fossil sources in Beijing (75 ± 10%), especially during haze periods. The fossil vs. non-fossil contributions to OC and EC did not change drastically during haze events in both Xi'an
- 25 and Beijing. However, compared to clean periods, the contribution of coal combustion to EC during haze periods increased in Xi'an and decreased in Beijing. During clean periods, primary OC from biomass burning and fossil sources constituted ~70% of OC in Xi'an and ~53% of OC in Beijing. From clean to haze periods, the contribution of SOC to total OC increased in Xi'an, but decreased in Beijing, suggesting that contribution of secondary organic aerosol formation to increased OC during haze periods was more efficient in Xi'an than in Beijing. In Beijing, the high SOC fraction in total OC during clean periods
- 30 was mainly due to elevated contribution from non-fossil SOC. In Xi'an, a slight day-night difference was observed during the clean period, with enhanced fossil contributions to OC and EC during the day. This day-night difference was negligible during severe haze periods, likely due to enhanced accumulation of pollutants under stagnant weather conditions.-

#### **1** Introduction

35 Severe haze pollution with high PM<sub>2.5</sub> (i.e., particulate matter with aerodynamic diameter ≤ 2.5 µm) concentrations and reduced visibility occurs frequently during winter in China (An et al., 2019). Filed Field measurements show that carbonaceous aerosol contributes a significant fraction of PM<sub>2.5</sub> loading during severe haze events in China (Huang et al., 2014; Elser et al., 2016; Liu et al., 2016). Therefore, a better understanding of the sources and atmospheric processes of carbonaceous aerosols is needed for mitigating haze pollution. Many previous studies focus solely on Beijing, the capital of China. However, studies 40 on other megacities are also needed for comparison as well as for a more comprehensive understanding of haze pollution in

China.

45

Carbonaceous aerosol constituents are separated into elemental carbon (EC) and organic carbon (OC), fractions differing in their thermal refractiveness with EC being thermally refractory and OC weakly refractory (Pöschl, 2003, 2005; Petzold et al., 2013). EC is emitted as primary particles from incomplete combustion sources (i.e., biomass burning and fossil fuel combustion). Unlike EC, OC can either be emitted as primary OC (POC) from combustion sources and non-combustion sources (e.g., biogenic emissions) or formed in the atmosphere as secondary OC (SOC) via the reaction of gas precursors (Hallquist et al., 2009; Jimenez et al., 2009). The sources and abundance of different carbon fractions in carbonaceous aerosols vary considerably in different Chinese cities, as a result of complex interplay between meteorology, local and regional

emissions sources, and atmospheric processes (Zhang et al., 2008; Cui et al., 2015; Tie et al., 2017; An et al., 2019). Therefore,

- 50 quantification the sources of carbonaceous aerosol in China is a challenging task.
  - Radiocarbon (<sup>14</sup>C) analysis of carbonaceous aerosols is the most direct and effective method to distinguish their main sources, exploiting the fact that OC and EC of fossil origins (i.e., vehicle emissions, coal combustion) do not contain <sup>14</sup>C (Heal, 2014; Cao et al., 2017; <u>Dusek et al., 2013</u>). <sup>14</sup>C analysis of OC and EC separately provide a clear-cut division of carbonaceous aerosols into four major fractions: fossil OC, non-fossil OC (e.g., OC from biomass burning, biogenic emissions and cooking),
- 55 fossil EC and biomass-burning EC (e.g., Gustafsson et al., 2009; Szidat et al., 2009; Zotter et al., 2014; Dusek et al., 2017; Ni et al., 2018, 2019a). For example, Liu et al. (2014) demonstrated that fossil sources including coal burning and vehicle emissions dominated EC during winter haze events in Guangzhou, China. Zhang et al. (2015) showed that the elevated carbonaceous aerosols during the severe haze event in January 2013 in China were by a large extent driven by SOC from both fossil and non-fossil precursors. In addition, the analysis of the <sup>13</sup>C/<sup>12</sup>C ratio can refine <sup>14</sup>C source apportionment, because
- 60 coal combustion and vehicle emissions have different <sup>13</sup>C source signatures although they both completely depleted in <sup>14</sup>C (e.g., Andersson et al., 2015; Li et al., 2016; Winiger et al., 2016, 2017; Fang et al., 2017, 2018; Ni et al., 2018).

A critical question for effective haze mitigation is whether carbonaceous aerosols in different Chinese cities have similar characteristics during haze events. However, there are not many studies highlighting the differences in sources of primary and secondary carbonaceous aerosols between cities, especially for studies employing the analysis of <sup>14</sup>C or the stable isotope <sup>13</sup>C (a.g., Zhang et al., 2015; Andersson et al., 2015; Liu et al., 2016). In this work, we compare the severe haze events reaching

65 (e.g., Zhang et al., 2015; Andersson et al., 2015; Liu et al., 2016). In this work, we compare the severe haze events reaching

"red alarm" level (i.e., the highest air-quality warning level in China) in two Chinese megacities (Beijing and Xi'an) during December 2016 and January 2017. We present measurements of dual carbon isotopes (i.e., <sup>14</sup>C and the stable carbon isotope <sup>13</sup>C) for EC and OC. The sources of carbonaceous aerosols are elucidated and compared between haze and clean periods in Beijing and Xi'an, with the main objectives: (1) quantitative understanding of the difference in EC contribution from burning

70

of biomass, coal and liquid fossil fuel (i.e., vehicle emissions) under different pollution conditions; and (2) constraint on the sources of both primary and secondary OC. Furthermore, the comparison of day-time and night-time results in Xi'an yields insight into diurnal variation in sources of carbonaceous aerosols.

#### 2 Methods

#### 2.1 Aerosol collection

- To collect PM<sub>25</sub> samples, high-volume aerosol samplers (flow rate =  $1.0 \text{ m}^3 \text{ min}^{-1}$ ; TE-6070 MFC, Tisch Inc., Cleveland, OH, 75 USA) were used at an urban background site in Xi'an and Beijing (see Table S1 for details about the sampling sites). Xi'an is the largest city in northwestern China, with over 8.8 million residents and 2.5 million vehicles in 2016 (Xi'an Municipal Bureau of Statistics and NBS Survey Office in Xi'an, 2017). Surrounded by Oinling Mountains to the south and the Loess Plateau to the north, days with low wind speed occur frequently in Xi'an, promoting the accumulation of air pollutants. Xi'an is now 80 facing increased serious air quality issues due to the rapid increase of motor vehicles and energy consumption in the past two decades. Besides residential coal combustion, biomass burning is also a major emission source in Xi'an and its surrounding areas (i.e., Guanzhong Plain) for heating and cooking especially in winter (Zhang et al., 2014; Xu et al., 2016). Beijing, the capital of China, is a megacity with over 21 million residents and 5.7 million vehicles in 2016 (Beijing Municipal Bureau of Statistics and NBS Survey Office in Beijing, 2017). Beijing is located in the Beijing-Tianjin-Hebei region, the most economically developed region in North China. However, the rapid economic growth and urbanization associated with heavy 85 coal consumption and rapid increase usage of vehicles lead to the poor air quality in Beijing. Besides local emissions, regional transport of pollutants between neighboring cities also contributes to air pollution in Beijing (Zheng et al., 2015; An et al., 2019). The 12 h integrated (daytime: 8:00 a.m. to 8:00 p.m., local standard time, LST; nighttime: 8:00 p.m. to 8:00 a.m. the following day)  $PM_{2.5}$  was sampled on pre-combusted quartz filters (8 in  $\times$  10 in; QM-A, Whatman Inc., Clifton, NJ, USA) in Xi'an from 1 January 2017 to 10 January 2017. In Beijing, the 24 h integrated (10:00 a.m. to 10:00 a.m. the following day) 90
- PM<sub>2.5</sub> was collected from 2 December 2016 to 7 January 2017. Field blanks were collected by exposing filters to ambient air for 15 min. Immediately after collection, the filters were transferred into pre-combusted aluminum foils and stored at -18 °C.

#### 2.2 Concentration measurements of OC and EC

IMPROVE A protocol (Chow et al., 2007) was implemented on a carbon analyzer (DRI Model 2001, Atmoslytic Inc., USA) for measurements of carbon concentrations. The relative standard deviations for the replicate analyses were smaller than 10 % 95 for OC and EC. OC mass was corrected for field blanks (0.4 µg cm<sup>-2</sup>). EC was too small to be detected on field blanks. Acidification to remove potential interferences from carbonates is not necessary, because carbonate carbon in PM<sub>2.5</sub> samples is found to be negligible, compared to the relatively larger OC and EC amounts for both mass determination and carbon isotopic analysis (Supplement S1).

#### 100 2.3 Analysis of carbon isotope

Six samples from haze and clean days were selected per sampling site for carbon isotope analysis (Tables S1 and S2, Fig. S1). We define clean days at each site as  $PM_{2.5} <$  median concentration in the winter heating season from 15 November 2016 to 15 March 2017. In Xi'an, there were 4 composite samples (2 daytime + 2 nighttime) from haze days, and 2 composite samples (1 daytime + 1 nighttime) from clean days. In Beijing, five 24 h samples were selected from haze days, and 1 composite sample

from two clean days. Each composite sample consists of 2 12h (for Xi'an) or 24 h (for Beijing) filter pieces with similar PM<sub>2.5</sub>
 loadings that agree within 20 % (Fig. S1).

### 2.3.1 Stable isotope <sup>13</sup>C

Filter samples were placed in a quartz tube with CuO grains. The tube was subsequently evacuated and sealed before heating for 3h at 375 °C to remove OC. Then the EC was extracted by heating the remaining carbon for 5 h at 850 °C. The  ${}^{13}C/{}^{12}C$ 

110 ratio of EC was measured by an isotope mass spectrometer (Finnigan MAT-251; Bremen, Germany) and expressed in the delta notation:

$$\delta^{13}C = \left[\frac{\binom{13}{C}\binom{12}{C}_{sample}}{\binom{13}{C}\binom{12}{C}_{V-PDB}} - 1\right].$$
(1)

 $\delta^{13}$ C values are usually reported in per mil (‰). ( $^{13}$ C/ $^{12}$ C)<sub>V-PDB</sub> is the  $^{13}$ C/ $^{12}$ C ratio of the international standard Vienna Pee Dee Belemnite (V-PDB). A well-characterized standard was measured every working day. Duplicate analysis of  $\delta^{13}$ C of EC showed an analytical precision better than  $\pm 0.3$ ‰. This method was detailed in Ni et al., (2019b), where impacts of potential charred OC on the isolated EC were evaluated using an isotope-mass-balance based sensitivity analysis. We concluded that the expected differences in  $\delta^{13}$ C<sub>EC</sub> is smaller than 1‰ under the assumption that the fraction of charred OC in the isolated EC is at most 20%.

2.3.2 Radiocarbon

115

- 120 OC and EC in PM<sub>2.5</sub> samples were converted to CO<sub>2</sub> using an aerosol combustion system (ACS; Dusek et al., 2014)<sub>25</sub>. The ACS has been evaluated in two intercomparison studies (Szidat et al., 2013; Zenker et al., 2017). The isolated CO<sub>2</sub> was subsequently reduced to graphite (de Rooij et al., 2010) before <sup>14</sup>C measurements can be conducted with the accelerator mass spectrometer (AMS) at CIO (van der Plicht et al., 2000). The temperature protocol for OC and EC combustion has been detailed in <del>Zenker</del> et al. (2017) and Ni et al. (2018), and is summarized in Fig. S2. To remove possible interfering gas (e.g., NO<sub>x</sub>, halogen and
- 125 water vapor) from CO<sub>2</sub>, a reduction oven filled with copper grains and silver, a dry ice-ethanol bath and a flask filled with phosphorus pentoxide are installed on the ACS.

Fraction modern ( $F^{14}C$ ) is used to report the <sup>14</sup>C data (Reimer et al., 2004).  $F^{14}C$  relates the <sup>14</sup>C/<sup>12</sup>C ratio of a sample to the ratio of the unperturbed atmosphere in the reference year 1950:

$$F^{14}C = \frac{\binom{{}^{14}C/{}^{12}C}_{\text{sample},[-25]}}{\binom{{}^{14}C/{}^{12}C}_{_{1950},[-25]}}.$$
(2)

Both ratios are normalized to  $\delta^{13}$ C of -25‰ to remove the effect of isotope fractionation. Practically,  $({}^{14}C/{}^{12}C)_{1950, [-25]}$  equals to the  ${}^{14}C/{}^{12}$ C ratio of an oxalic acid standard (OXII) multiplied by a factor of 0.7459. Contamination during graphitization and AMS measurements was quantified from the measured F<sup>14</sup>C of standards (OXII with known F<sup>14</sup>C of 1.3407 and Rommenhöller with F<sup>14</sup>C=0) processed in the same way as samples. The resulting estimated dead and modern contamination were used to correct the <sup>14</sup>C data according to Santos et al. (2007). The reliability of data correction was further verified by measuring two secondary standards (i.e., IAEA-C7 and-C8) on the same wheel of samples. The measured values of IAEA-C7 (0.495 ± 0.008) and IAEA-C8 (0.154 ± 0.007) agree with their respective consensus value (0.4953 ± 0.0012 and 0.1503 ±

0.0017) within uncertainties.

#### 2.4 Source apportionment methods

F<sup>14</sup>C is larger than the fraction of non-fossil carbon (i.e., *f*<sub>nf</sub>(OC) for OC, *f*<sub>bb</sub>(EC) for EC) due to the large release of <sup>14</sup>C into
the atmosphere from the nuclear bomb tests in 1960s. To eliminate this effect, F<sup>14</sup>C is divided by F<sup>14</sup>C of non-fossil sources (F<sup>14</sup>C<sub>nf</sub>). F<sup>14</sup>C<sub>nf</sub> is estimated as 1.09 ± 0.05 for OC and 1.10 ± 0.05 for EC (see details in Ni et al., 2019b), using a tree growth model and the contemporary atmospheric <sup>14</sup>CO<sub>2</sub> over the past years (Lewis et al., 2004; Mohn et al., 2008; Levin et al., 2010), with the assumption that biomass-burning OC and biogenic OC contribute to 85% and 15% of total OC, respectively. Once *f*<sub>nf</sub>(OC) and *f*<sub>nf</sub>(EC) are known, carbon concentrations can be apportioned into EC and OC from non-fossil sources (EC<sub>bb</sub>, OC<sub>nf</sub>)
and fossil sources (EC<sub>fossil</sub>, OC<sub>fossil</sub>) (Eq. 3–6 in Table 1). OC<sub>nf</sub> and OC<sub>fossil</sub> are further divided into POC from biomass burning (POC<sub>bb</sub>), other non-fossil OC (OC<sub>o,nf</sub>) (Eq. 7–8), primary and secondary fossil OC (POC<sub>fossil</sub> and SOC<sub>fossil</sub>, respectively; Eq. 9–10). POC<sub>bb</sub> and POC<sub>fossil</sub> are estimated using EC as a tracer of primary emissions (i.e., the EC tracer method; Turpin and Huntzicker, 1995). Based on OC<sub>o,nf</sub> and SOC<sub>fossil</sub>, total SOC and the fraction of fossil carbon in SOC (*f*<sub>fossil</sub>(SOC)) are estimated using Eq. (11–12). OC<sub>o,nf</sub> mainly includes SOC of non-fossil origins (SOC<sub>nf</sub>), primary biogenic OC and cooking OC. OC<sub>o,nf</sub> is approximately SOC<sub>nf</sub>, as contributions of primary biogenic sources and cooking to OC<sub>o,nf</sub> are likely small (Hu et al., 2010; Guo

et al., 2012). If cooking is prominent,  $OC_{o,nf}$  is an overestimate of  $SOC_{nf}$ . To estimate the uncertainties of the source apportionment results, a Monte Carlo simulation (n=10000) using Eq. (3–12) was carried out as described in Supplement S2S3. The <sup>14</sup>C source apportionment results are presented in Tables S3 and S4.

The dual carbon isotope signatures of EC were used in a Bayesian Markov chain Monte Carlo (MCMC) scheme (Andersson, 2011), to conduct the mass-balance three source apportionment of EC (e.g., Andersson et al., 2015; Li et al., 2016; Winiger et

al., 2016, 2017; Fang et al., 2017, 2018). That is, the F<sup>14</sup>C and  $\delta^{13}$ C of ambient EC (F<sup>14</sup>C<sub>(EC)</sub> and  $\delta^{13}$ C<sub>EC</sub>) can be explained by burning of biomass (<sub>bb</sub>), coal (<sub>coal</sub>) and liquid fossil fuel (<sub>liq.fossil</sub>; i.e., vehicle emissions):

$$\begin{pmatrix} F^{14}C_{(EC)} \\ \delta^{13}C_{EC} \\ 1 \end{pmatrix} = \begin{pmatrix} F^{14}C_{nf} & F^{14}C_{coal} & F^{14}C_{liq.fossil} \\ \delta^{13}C_{bb} & \delta^{13}C_{coal} & \delta^{13}C_{liq.fossil} \\ 1 & 1 & 1 \end{pmatrix} \begin{pmatrix} f_{bb} \\ f_{coal} \\ f_{liq.fossil} \end{pmatrix}$$
(13)

 $F^{14}C_{coal}$  and  $F^{14}C_{liq,fossil}$  equal to zero since coal and liquid fossil fuel do not contain  ${}^{14}C. \delta^{13}C_{bb}, \delta^{13}C_{coal}$  and  $\delta^{13}C_{liq,fossil}$  are  $\delta^{13}C$ signatures for EC from the three sources. Their values were established as  $\delta^{13}C_{bb}$  (-26.7 ± 1.8 ‰ for C3 plants, and -16.4 ± 1.4 ‰ for corn stalk; mean ± SD),  $\delta^{13}C_{coal}$  (-23.4 ± 1.3 ‰) and  $\delta^{13}C_{liq,fossil}$  (-25.5 ± 1.3 ‰), based on critical evaluations of literature studies (Andersson et al., 2015; Ni et al., 2018; and references therein). Uncertainties in F<sup>14</sup>C and  $\delta^{13}C$  source signatures and the measured  $F^{14}C_{(EC)}$  and  $\delta^{13}C_{EC}$  are considered in the MCMC technique (Parnell et al., 2010, 2013). MCMC outputs are the posterior PDFs for  $f_{bb}$ ,  $f_{coal}$  and  $f_{liq,fossil}$  (i.e., the relative contribution of each source to EC). The median and interquartile range (25th–75th percentile) are used as the best estimate and the uncertainties, respectively.

#### **3** Results and discussion

#### 3.1 Fossil and non-fossil contributions to EC and OC

During the measurement periods, the highest daily mass concentrations of  $PM_{2.5}$  in Xi'an (~250–420 µg m<sup>-3</sup>) and Beijing (~210–360 µg m<sup>-3</sup>; Fig. S1) were 10–17 and 8–14 times higher than the standard of World Health Organization (25 µg m<sup>-3</sup>;

- WHO, 2006), respectively. Using radiocarbon measurements, we investigated the sources of carbonaceous aerosols in PM<sub>2.5</sub> in both cities during several haze periods, and compared them to clean periods, with PM<sub>2.5</sub> concentrations below 100 µg m<sup>-3</sup> in Xi'an and below 20 µg m<sup>-3</sup> in Beijing. In Xi'an, even during clean periods we defined here, daily PM<sub>2.5</sub> concentrations were higher than the Chinese pollution standard of 75 µg m<sup>-3</sup>, reflecting severe air quality problems. PM<sub>2.5</sub>, OC and EC concentrations during haze periods were > 2 times higher in Xi'an and > 5 times higher in Beijing than those during clean periods, respectively. OC/EC ratios in Xi'an slightly decreased from ~4 during haze periods to ~3 during clean periods, while
- OC/EC ratios in Beijing were lower during haze periods (~3) than clean periods (~4). This reflects different sources and formation mechanisms of haze pollution in the two cities. In Xi'an, we collected day and night PM<sub>2.5</sub> samples. No consistent day-night variations in concentrations of PM<sub>2.5</sub>, OC and EC (Figs. 1 and S1) were observed. This is resulted from diurnal cycle of human activities (e.g., traffic, usage of biomass and coal for heating or cooking) and the development of planetary boundary layer height which controls the vertical mixing and dilution of pollutants.
- 180 layer height which controls the vertical mixing and dilution of pollutants.

Radiocarbon (<sup>14</sup>C) in EC and OC was measured to distinguish their fossil (mainly coal burning and traffic emissions) and nonfossil sources (mainly biomass burning). The most important contributor to EC was fossil fuel combustion, both in Xi'an and Beijing, contributing  $73 \pm 2\%$  in Xi'an and  $80 \pm 3\%$  in Beijing. The remaining EC arose from biomass burning ( $27 \pm 2\%$  in Xi'an and  $20 \pm 3\%$  in Beijing; Fig. 1). In Xi'an, the fraction of biomass-burning EC in total EC ( $f_{bb}(EC)$ ) was largely constant

- during haze and clean periods (range: 25%–29%), regardless of the wide concentration range of EC from biomass burning 185  $(EC_{bb}, 1.8-6.4 \ \mu g \ m^{-3})$  and fossil fuel combustion  $(EC_{fossil}, 4.3-18 \ \mu g \ m^{-3})$ . This suggests that the increase in  $EC_{fossil}$  and  $EC_{bb}$ concentrations during haze periods in Xi'an is likely caused by the enhanced emissions from both fossil fuel and biomass burning by a similar factor and due to meteorological conditions favoring the accumulation of particulate air pollutions.  $f_{bb}(EC)$ values in Beijing  $(20 \pm 3\%)$  with a range of 17%–24%) were consistently smaller than those in Xi'an (range: 25%–29%), 190 showing that fossil sources contribute more strongly to EC in Beijing. Moreover, during haze periods in Beijing,  $f_{\rm bb}(\rm EC)$
- increased with increasing total EC concentrations (Fig. 2).

In Xi'an, OC concentrations from non-fossil sources averaged  $29 \pm 16 \,\mu g \,m^{-3}$  (OC<sub>nf</sub>; range: 9–49  $\mu g \,m^{-3}$ ), slightly higher than those from fossil sources (OC<sub>fossil</sub>;  $24 \pm 13 \ \mu g \ m^{-3}$ ; range: 8–40  $\mu g \ m^{-3}$ ) at 95% confidence level (paired t test, p-value = 0.01). However, in Beijing,  $OC_{pf}$  (12 ± 5 µg m<sup>-3</sup>; 3–19 µg m<sup>-3</sup>) was significantly lower than  $OC_{fossil}$  (24 ± 10 µg m<sup>-3</sup>; 4–33 µg m<sup>-3</sup>) (p-

value = 0.001). Consequently, the relative contribution of OC<sub>nf</sub> to total OC ( $f_{nf}(OC)$ ) was much higher in Xi'an (average 54 ± 195 2 %) than in Beijing ( $34 \pm 3\%$ ).  $f_{nf}(OC)$  in both cities was considerably higher than the corresponding  $f_{bb}(EC)$  for all samples (Fig. 1). The main reason for larger  $f_{\rm nf}(OC)$  than  $f_{\rm bb}(EC)$  is that primary OC/EC ratios from biomass burning emissions are higher than those from fossil sources. So even though biomass burning contributes a small portion of EC, its contribution to primary OC will be much higher. In addition, other non-combustion sources (e.g., biogenic emissions, cooking fumes) and 200 secondary formation contribute only to OC, but not to EC.

In this study, the  $f_{\text{fossil}}(\text{EC})$  values in Xi'an during winter 2016/2017 are comparable with those previously measured during winter 2015/2016 and winter 2008/2009 (Ni et al., 2018, 2019b), as illustrated in Fig. 3b, pointing to relative constant contribution of fossil fuel combustion vs. biomass burning to EC in Xi'an over the past decade. As shown in Fig. 3b, the  $f_{\text{tossil}}(\text{EC})$  values in Beijing during winter 2016/2017 agree with the values reported at an urban site of Beijing in January 2014

- 205 (Fang et al., 2017). A slightly higher  $f_{\text{fossil}}(\text{EC})$  in urban Beijing was observed during February 2010 (Chen et al., 2013). Despite the slight variation of f<sub>rossil</sub>(EC) over time, f<sub>rossil</sub>(EC) in Beijing is generally higher than that in Xi'an (Fig. 3b). The presented overall average  $f_{\text{fossil}}(\text{OC})$  for winter 2016/2017 in Beijing (66 ± 3%) was higher than that in Xi'an (46 ± 3%), consistent with previously reported  $f_{\text{fossil}}(OC)$  in Beijing and Xi'an (Zhang et al., 2015; Ni et al., 2019a). Lower  $f_{\text{fossil}}(OC)$  values in winter were reported for Chongqing (24%), and higher  $f_{\text{fossil}}(OC)$  was observed in Taiyuan (71%) during winter 2013/2014 (Ni et al.,
- 210 2019a). The comparison of  $f_{\text{fossil}}(\text{EC})$  and  $f_{\text{fossil}}(\text{OC})$  in different Chinese cities indicates that the relative importance of fossil sources in carbonaceous aerosols vary spatially, and can change over the years. In Xi'an, clean periods showed a slight daynight difference with increased contributions of fossil sources to EC and OC during the day. During haze periods, especially the 2<sup>nd</sup> haze event (XH\_day2, XH\_night2), this day-night difference disappeared, which suggests a long residence time of the pollution particles in the urban atmosphere during haze events.
- 215 Overall, our <sup>14</sup>C data show that fossil sources contribute more strongly to EC and OC in Beijing than in Xi'an, which is consistent with previous observations. Both in Beijing and Xi'an the fossil vs. non-fossil contributions to EC and OC did not

change drastically during haze and clean periods. In Xi'an, a slight day-night difference was observed during clean periods, but disappeared during haze periods, suggesting a large accumulation of particles.

#### 3.2 Fossil EC apportioned by stable carbon isotopes: coal vs. liquid fossil fuel

Besides F<sup>14</sup>C, the  $\delta^{13}$ C of EC adds additional dimension where fossil EC can be distinguished into EC from burning of coal and of liquid fossil fuel (i.e., vehicle emissions). Considerable geographical differences in  $\delta^{13}C_{EC}$  signatures were observed, with more depleted values in Xi'an (-25.1 ± 0.5%; -25.6% to -24.4 ‰) relative to those in Beijing (-24.1 ± 0.4‰; -24.4‰ to -23.4‰; Fig. 3). The Xi'an signatures are consistent with the signature of liquid fossil fuel combustion ( $\delta^{13}C_{liq,fossil} = -25.5 \pm 1.3\%$ ; Sect. 2.4), whereas the more enriched values in Beijing indicate the influence of coal combustion ( $\delta^{13}C_{coal} = -23.4 \pm 1.3\%$ ).

In both Xi'an and Beijing, moderate differences exist in  $\delta^{13}C_{EC}$  between clean and haze days, pointing to a shift in combustion sources. In Xi'an,  $\delta^{13}C_{EC}$  during clean periods (~-25.5‰) was slightly depleted compared to that during haze periods (-25.0‰ to -24.4‰), whereas Beijing exhibited more enriched  $\delta^{13}C_{EC}$  during clean periods (-23.4‰) than during haze periods (-24.4‰ to -24.1‰). This suggests a moderate increase in coal combustion contribution to EC in Xi'an during haze days and a decrease

230 in Beijing. In Xi'an, no strong day-night difference in  $\delta^{13}C_{EC}$  was observed, with the largest absolute differences of 0.5‰ between XH\_day1 and XH\_night1. The day-night differences are small relative to the uncertainties of the potential sources, for example, the endmember range for coal combustion is more uncertain (± 1.3‰). The small day-night differences in  $\delta^{13}C_{EC}$  reflect well-mixed EC emissions.

The Bayesian MCMC model takes into account the uncertainties of the  $\delta^{13}$ C and  $F^{14}$ C endmembers and statistically apportions

- EC into the fraction of biomass burning ( $f_{bb}$ ), coal combustion ( $f_{coal}$ ) and liquid fossil fuel combustion ( $f_{liq,fossil}$ ). The MCMCderived  $f_{bb}$  is in principle the same as the <sup>14</sup>C-based  $f_{bb}$ (EC) (Fig. S3). The MCMC results (Fig. 4) show that there were no strong day-night differences in EC sources during haze and clean periods in Xi'an. Liquid fossil fuel combustion was the most important contributor to EC in Xi'an, with increased contribution during clean periods. In Beijing, coal combustion was the dominant source of EC, with the relative contribution ranging from 48% (median; 31%–61%, interquartile range) during haze
- 240 periods to 61% (45%–71%) during clean periods.  $f_{bb}$  was fairly constant between haze and clean periods with respect to  $f_{coal}$ and  $f_{iiq.fossil}$  for all samples. In Xi'an,  $f_{bb}$  was comparable to  $f_{coal}$  during haze days, and larger than  $f_{coal}$  during clean days. In Beijing, biomass-burning EC was the smallest fraction in total EC, with smaller  $f_{bb}$  than  $f_{coal}$  during both haze and clean days. Concentrations of total EC increased by 2 times from clean days (~7.4 µg m<sup>-3</sup>) to haze days (18.0 µg m<sup>-3</sup>) in Xian, and 8 times in Beijing (1.6 µg m<sup>-3</sup> to 13.5 µg m<sup>-3</sup>). The increased EC concentrations during haze periods in Xi'an were attributed to liquid
- fossil fuel combustion (43%), coal combustion (29%) and biomass burning (28%). However, in Beijing, coal combustion contributed most of the increased concentrations of EC (45%), followed by burning of liquid fossil fuel (33%) and biomass (22%).

In summary, complementing <sup>14</sup>C with  $\delta^{13}$ C allows for quantitative constraints on EC sources: EC was dominated by liquid fossil fuel combustion (i.e., vehicle emissions) in Xi'an and by coal burning in Beijing, especially during clean periods. In

- Xi'an, no strong day-night differences in EC sources were observed during haze and clean periods. Compared with earlier observations in Xi'an (Fig. 3b), we found that the δ<sup>13</sup>C<sub>EC</sub> values in January 2017 from this study are comparable with wintertime δ<sup>13</sup>C<sub>EC</sub> in 2015/2016 (Ni et al., 2019b), but much more depleted than wintertime δ<sup>13</sup>C<sub>EC</sub> in 2008/2009 (Ni et al., 2018) and January 2003 (Cao et al., 2011). This suggests that fossil sources of EC in Xi'an have changed in the past decade, with decreasing relative contribution from coal combustion. This is in line with recent changes in energy use, and the decreasing enrichment factors of As and Pb (e.g., indicators of coal combustion) in Xi'an, as documented in recent studies (Xu et al., 2016). As shown in Fig. 3b, in Beijing, variations in δ<sup>13</sup>C<sub>EC</sub> from January 2003 (Cao et al., 2011) to January 2017 (this study) are much narrower than those in Xi'an, indicating that EC combustion sources did not change significantly throughout the years in Beijing. Our δ<sup>13</sup>C<sub>EC</sub> values overlap with those in January 2014 (Fang et al., 2017) and fall into the range of reported δ<sup>13</sup>C<sub>EC</sub> values in urban Beijing (Cao et al., 2011; Chen et al., 2013) and the regional receptor site of Beijing (Andersson et al., 2015; Fang et al., 2017).
  - 3.3 Primary and Secondary secondary OC

As explained in Sect. 2.4,  $OC_{nf}$  and  $OC_{fossil}$  are apportioned into primary (POC<sub>bb</sub>, POC<sub>fossil</sub>) and secondary OC (OC<sub>o,nf</sub>, SOC<sub>fossil</sub>; Fig. 5). The large error bars of their concentrations reflect the large uncertainties in  $r_{bb}$  and  $r_{fossil}$ . It should be noticed that  $OC_{o,nf}$  is used as an approximation of SOC<sub>nf</sub>, or can be regarded as an upper limit of SOC<sub>nf</sub> if cooking is a prominent OC source.

- In Xi'an, both ratios of  $OC_{o,nf}/POC_{bb}$  and  $SOC_{fossil}/POC_{fossil}$  increased during haze periods (Fig. 5a).  $OC_{o,nf}/POC_{bb}$  ratio increased by 2.5 times from 0.33–0.46 during clean periods to 0.86–1.1 during haze periods, in contrast to  $SOC_{fossil}/POC_{fossil}$ increased by 1.5 times from 0.46–0.50 to 0.62–0.78. This underlines that haze episodes in Xi'an were mainly caused by additional SOC formation, with larger contribution from non-fossil sources than fossil sources. As shown in Fig. 5b, the contribution of SOC (i.e.,  $SOC \cong OC_{o,nf} + SOC_{fossil}$ ) to OC increased from clean periods (28%–32%) to haze periods (44%–
- 270 48%), mainly resulted from increased contribution of  $OC_{o,nf}$  to total OC (i.e., from 14%–16% to 26%–29%). In Xi'an, the daynight difference was larger during clean periods with less SOC at night for both absolute concentration and relative contribution to total OC (Figs. 5b, 5c).

In contrast, Beijing had the opposite variation trends of  $OC_{o,nf}/POC_{bb}$  and  $SOC_{fossil}/POC_{fossil}$  from clean to haze periods.  $OC_{o,nf}/POC_{bb}$  ratios during clean periods (1.3) were on average five times higher than those during haze periods (0.18–0.33),

and  $SOC_{fossil}/POC_{fossil}$  ratios during clean periods (0.71) were slightly higher than those during haze periods (0.41–0.64). This suggests that in Beijing the increased OC concentrations during haze periods were mainly derived from elevated concentrations of  $POC_{bb}$  and  $POC_{fossil}$ . As shown in Fig. 5b, high SOC contribution to total OC was observed during clean periods, mainly due to elevated contribution from  $OC_{o,nf}$ . In winter, the The  $OC_{o,nf}$  is not likely attributed to biogenic OC, because the biogenic emissions are very low in winter. As a result, the elevated contribution from  $OC_{o,nf}$  to OC during clean periods in Beijing could

280 be attributed to regional sources. During clean periods, concentrations of OC and  $OC_{o,nf}$  are small, and the measured carbon concentrations can reflect regional sources, which are dominated by secondary sources due to long-range transport. It could also be that contribution of cooking OC to  $OC_{o,nf}$  can be noticeable during clean conditions.

The fossil fraction of the total SOC can be defined as  $f_{\text{fossil}}(\text{SOC}) = \text{SOC}_{\text{fossil}}/\text{SOC}$ . In Xi'an around half of SOC was derived from fossil sources ( $f_{\text{fossil}}(\text{SOC}) = 44 \pm 6\%$ ), whereas  $f_{\text{fossil}}(\text{SOC}) = 75 \pm 10\%$  in Beijing. Using a similar approach with this study, Zhang et al. (2015) also found that Beijing had higher  $f_{\text{fossil}}(\text{SOC})$  (48%–63%) than in Xi'an (30%–35%) during the January 2013 severe haze events. These findings suggest the important contribution of fossil sources to SOC in Beijing and non-fossil sources in Xi'an.  $f_{\text{fossil}}$  (SOC) in Beijing increased during haze periods, whereas the opposite trend was found in Xi'an (Fig. 6). During haze periods in Beijing,  $f_{\text{fossil}}$ (SOC) overlapped with  $f_{\text{fossil}}$ (EC), and was clearly higher than  $f_{\text{fossil}}$ (OC).

Together these results reveal the differences in primary and secondary OC in two Chinese megacities. The contribution of SOC to total OC increased from clean to haze periods in Xi'an. In contrary, SOC/OC ratios increased from haze to clean periods in Beijing, mainly due to increased SOC from non-fossil sources. SOC was dominated by fossil sources in Beijing but by non-fossil sources in Xi'an, especially during haze periods. In Xi'an, the day-night difference was larger during clean periods with less SOC at night.

#### 3.4 Differences between the fractions of non-fossil carbon in OC and EC

295 The differences between  $f_{nf}(OC)$  and  $f_{bb}(EC)$  were smaller in Beijing, ranging from 11% to 20%, compared to 25%–29% in Xi'an. To better understand what governs the differences, we express  $f_{nf}(OC)$  in terms of fossil to biomass burning ratio in EC and primary OC/EC emissions ratios. Starting from the formulas of  $f_{bb}(EC)$  and  $f_{nf}(OC)$ :

$$f_{bb}(EC) = \frac{EC_{bb}}{EC_{bb} + EC_{fossil}} = \frac{1}{1 + \frac{EC_{fossil}}{EC_{bb}}}$$
(14)

$$f_{\rm nf}(\rm OC) = \frac{OC_{\rm nf}}{OC_{\rm nf} + OC_{\rm fossil}} = \frac{POC_{\rm bb} + OC_{\rm o,nf}}{POC_{\rm bb} + OC_{\rm o,nf} + POC_{\rm fossil} + SOC_{\rm fossil}}$$
(15)

300 We find that

$$f_{nf}(OC) = \frac{1}{1 + \frac{(1 + SOC_{fossil}/POC_{fossil})}{(1 + OC_{o,nf}/POC_{bb})} \times \frac{POC_{fossil}}{POC_{bb}}} = \frac{1}{1 + \frac{(1 + SOC_{fossil}/POC_{fossil})}{1 + \frac{(1 + SOC_{fossil}/POC_{fossil})}{(1 + OC_{o,nf}/POC_{bb})} \times \frac{r_{fossil}}{r_{bb}} \times \frac{EC_{fossil}}{EC_{bb}}}$$
(16)

where  $r_{\text{fossil}}$  is the weighted average of  $r_{\text{coal}}$  and  $r_{\text{vehicle.}}$ 

Comparing Eq. (14) with Eq. (16), we find that  $f_{nf}(OC)$  and  $f_{bb}(EC)$  would be equal if  $\frac{(1+SOC_{fossil}/POC_{fossil})}{(1+OC_{o,nf}/POC_{bb})} \times \frac{r_{fossil}}{r_{bb}} = 1$ . Since

 $r_{\text{fossil}}$  is usually smaller than  $r_{\text{bb}}$ ,  $f_{\text{nf}}(\text{OC})$  tends to be larger than  $f_{\text{bb}}(\text{EC})$ , assuming that SOC formation is comparable for fossil 305 or non-fossil sources (i.e.,  $\frac{(1+\text{SOC}_{\text{fossil}}/\text{POC}_{\text{fossil}})}{(1+\text{OC}_{\text{o,nf}}/\text{POC}_{\text{bb}})} \sim 1$ ). With smaller  $r_{\text{fossil}}$  than  $r_{\text{bb}}$  similar  $f_{\text{nf}}(\text{OC})$  and  $f_{\text{bb}}(\text{EC})$  can result from larger secondary formation from fossil sources than non-fossil sources (i.e.,  $\frac{(1+\text{SOC}_{\text{fossil}}/\text{POC}_{\text{fossil}})}{(1+\text{OC}_{\text{o,nf}}/\text{POC}_{\text{bb}})} > 1$ ). However, the fossil source coal combustion has a higher primary OC to EC ratio than vehicle emissions (i.e.,  $r_{\text{coal}}>r_{\text{vehicle}}$ ). Therefore, in a city where biomass burning and coal combustion are the dominant pollution sources,  $f_{\text{nf}}(\text{OC})$  and  $f_{\text{bb}}(\text{EC})$  will be more similar than in a city where the main sources are biomass burning and vehicle emissions.

- 310 Compared to Xi'an, Beijing had significantly smaller differences between  $f_{bb}(EC)$  and  $f_{nf}(OC)$  (Fig. 1), which was also observed in previous studies during the haze event in January 2013 (Zhang et al., 2015). Comparing Eq. (14) with Eq. (16), This-this suggests either strong contribution from coal combustion in Beijing or large secondary formation from fossil sources, or both. The stronger contribution of coal combustion to OC in Beijing than in Xi'an was a direct consequence of a larger proportion of coal combustion in EC in Beijing, as demonstrated by the Bayesian MCMC results of EC (Sect. 3.2). The latter was further
- 315 validated by the variation of SOC. The  $f_{\text{fossil}}(\text{SOC})$  in Beijing was higher than that in Xi'an, despite the variations between haze and clean periods (Sect. 3.3). By combining <sup>14</sup>C measurements with other state-of-art analytical techniques (e.g., aerosol mass spectrometry), Huang et al. (2014) also found that fossil OC was mostly secondary in nature in Beijing, and non-fossil SOC formation was dominant in Xi'an during a wintertime haze episode (i.e., Beijing had a larger  $f_{\text{fossil}}(\text{SOC})$  than Xi'an). However, atmospheric mechanisms responsible for the enhancement in fossil-derived SOA formation in Beijing remain unclear.
- 320 Furthermore, as shown in Fig. 1, unlike Xi'an where the differences between  $f_{nf}(OC)$  and  $f_{bb}(EC)$  were relatively constant for all samples, in Beijing the differences between  $f_{nf}(OC)$  and  $f_{bb}(EC)$  were smaller during haze periods than clean periods, caused by decreased  $f_{bb}(EC)$  and slightly increased  $f_{nf}(OC)$  during clean periods. This might indicate a higher relative contribution from coal combustion and/or fossil-dominated SOC during haze periods in Beijing. However, the Bayesian MCMC results of EC show the opposite, i.e., in Beijing the contribution of coal combustion to EC was lower during haze periods than during
- 325 clean periods (Sect. 3.2). Therefore, the only possible explanation is that, during haze periods in Beijing, SOC was dominated by fossil sources. This is validated by considerable higher significantly larger  $f_{\text{fossil}}(\text{SOC})$  during haze periods (76%–81%) than during clean periods (~55%; Sect. 3.3).

In conclusion, this section discusses the factors that governs the differences between  $f_{bb}(EC)$  and  $f_{nf}(OC)$ , and concludes that smaller differences suggest stronger contribution from coal combustion and/or larger secondary formation from fossil sources. This is further examined and validated by source apportionment results of EC and OC (Sect. 3.1–3.3).

#### **4** Conclusion

330

In this study the sources of carbonaceous aerosol were quantified using a dual-carbon isotopic approach for PM<sub>2.5</sub> samples collected in urban Xi'an and Beijing reaching "red alarm" level during December 2016 and January 2017. The <sup>14</sup>C results

showed that fossil sources dominated EC, contributing on average  $73 \pm 2$  % of EC in Xi'an and  $80 \pm 3$ % of EC in Beijing. The

remaining EC was attributed to biomass burning. In Xi'an, *f*<sub>bb</sub>(EC) was fairly constant during haze and clean periods, despite the wide <u>variation-range</u> of EC concentrations. However, in Beijing, *f*<sub>bb</sub>(EC) increased with increasing EC concentrations. Complementing <sup>14</sup>C with δ<sup>13</sup>C in a Bayesian MCMC approach allows for separation of fossil sources of EC into coal combustion and liquid fossil fuel combustion. The MCMC results in Xi'an suggest that liquid fossil fuel combustion contributed 44%–49% of EC during haze periods, and 54%–57% of EC during clean periods. In Beijing, coal combustion was the dominate fossil source of EC, with decreasing contribution to EC from clean periods (~61%) to haze periods (~48%).

<sup>14</sup>C measurements of OC showed that the contribution of non-fossil sources to OC was larger than that to EC, and was on average  $54 \pm 2$  % in Xi'an and  $34 \pm 3$ % in Beijing. The differences between non-fossil fraction in OC and EC were smaller in Beijing and larger in Xi'an. This suggests strong contribution from coal combustion or large secondary formation from fossil sources, or both in Beijing. In Xi'an, the fraction of SOC in total OC was larger during haze periods than during clean periods,

345 mainly due to increased SOC from non-fossil sources. Beijing showed the opposite trends with a larger fraction of SOC in total OC during clean periods than during haze periods, mainly due to elevated contribution from non-fossil SOC during clean periods.

SOC was dominated by non-fossil sources in Xi'an but by fossil sources in Beijing, especially during haze periods. The relative contribution of fossil sources to SOC ( $f_{\text{fossil}}(\text{SOC})$ ) was consistently higher in Beijing than in Xi'an. In Beijing,  $f_{\text{fossil}}(\text{SOC})$ 

- 350 was higher during haze periods (76%–81%) than during clean periods (55%), whereas an opposite trend was found in Xi'an, with  $f_{\text{fossil}}(\text{SOC})$  increasing from ~39%–43% during haze periods to ~52% during clean periods. In Xi'an, a slight day-night difference was found during clean periods, with increasing fossil contribution to OC and EC during the day and less SOC at night. During strong haze, this day-night difference was negligible, suggesting a large accumulation under stagnant weather conditions during the severe haze periods.
- 355 *Data availability.* Data used to support the findings in this study are archived at the Institute of Earth Environment, Chinese Academy of Sciences, and are available on request from the corresponding author.

Competing interests. The authors declare that they have no conflict of interest.

*Author contributions.* RJH and UD designed the study. Isotope measurements were made by HN, MMC, and JG. Data analysis and interpretation were made by HN, MMC, RJH, and UD. HN wrote the paper with contributions from all co-authors.

360 *Acknowledgments.* We acknowledge the financial support from the Gratama Foundation. We thank Marc Bleeker and Henk Been for their help with the <sup>14</sup>C measurements. Special thanks are also given to Anita Aerts-Bijma and Dicky van Zonneveld for their assistance in<sup>14</sup>C data correction at CIO.

*Financial support*. The authors acknowledge the project grants from the National Key Research and Development Program of China (no. 2017YFC0212701), the National Natural Science Foundation of China (nos. 41925015, 91644219, and 41877408),

365 the KNAW (no. 530-5CDP30), the Chinese Academy of Sciences (nos. ZDBS-LY-DQC001), and the Cross Innovative Team fund from the State Key Laboratory of Loess and Quaternary Geology (no. SKLLQGTD1801).

#### References

395

An, Z., Huang, R.-J., Zhang, R., Tie, X., Li, G., Cao, J., Zhou, W., Shi, Z., Han, Y., Gu, Z., and Ji, Y.: Severe haze in northern China: a synergy of anthropogenic emissions and atmospheric processes, P. Natl. Acad. Sci. USA, 116, 8657–8666, https://doi.org/10.1073/pnas.1900125116, 2019.

Andersson, A.: A systematic examination of a random sampling strategy for source apportionment calculations, Sci. Total Environ., 412, 232–238, 2011.

Andersson, A., Deng, J., Du, K., Zheng, M., Yan, C., Sköld, M., and Gustafsson, Ö.: Regionally-varying combustion sources of the January 2013 severe haze events over eastern China, Environ. Sci. Technol., 49, 2038–2043, 2015.

375 <u>Beijing Municipal Bureau of Statistics and NBS Survey Office in Beijing: Beijing Statistical Yearbook, China Statistics Press,</u> <u>Beijing, China, 2017 (in Chinese).</u>

Cao, F., Zhang, Y., Ren, L., Liu, J., Li, J., Zhang, G., Liu, D., Sun, Y., Wang, Z., Shi, Z., and Fu, P.: New insights into the sources and formation of carbonaceous aerosols in China: potential applications of dual-carbon isotopes, Natl. Sci. Rev., 4, 804–806, 2017.

380 Cao, J.-J., Chow, J. C., Tao, J., Lee, S. C., Watson, J. G., Ho, K. F., Wang, G. H., Zhu, C. S., and Han, Y. M.: Stable carbon isotopes in aerosols from Chinese cities: influence of fossil fuels, Atmos. Environ., 45, 1359–1363, 2011.

Chen, B., Andersson, A., Lee, M., Kirillova, E. N., Xiao, Q., Kruså, M., Shi, M., Hu, K., Lu, Z., Streets, D. G., Du, K., and Gustafsson, Ö.: Source forensics of black carbon aerosols from China, Environ. Sci. Technol., 47, 9102–9108, 2013.

Chow, J. C., Watson, J. G., Chen, L.-W. A., Chang, M. O., Robinson, N. F., Trimble, D., and Kohl, S.: The IMPROVE\_A
 temperature protocol for thermal/optical carbon analysis: maintaining consistency with a long-term database, J. Air Waste
 Manage., 57, 1014–1023, 2007.

Cui, H., Mao, P., Zhao, Y., Nielsen, C. P., and Zhang, J.: Patterns in atmospheric carbonaceous aerosols in China: emission estimates and observed concentrations, Atmos. Chem. Phys., 15, 8657–8678, https://doi.org/10.5194/acp-15-8657-2015, 2015.

de Rooij, M., van der Plicht, J., and Meijer, H.: Porous iron pellets for AMS <sup>14</sup>C analysis of small samples down to ultramicroscale size ( $10-25 \mu g C$ ), Nucl. Instrum. Meth. B, 268, 947–951, 2010.

Dusek, U., Ten Brink, H., Meijer, H., Kos, G., Mrozek, D., Röckmann, T., Holzinger, R., and Weijers, E.: The contribution of fossil sources to the organic aerosol in the Netherlands, Atmos. Environ., 74, 169–176, 2013.

Dusek, U., Monaco, M., Prokopiou, M., Gongriep, F., Hitzenberger, R., Meijer, H. A. J., and Röckmann, T.: Evaluation of a two-step thermal method for separating organic and elemental carbon for radiocarbon analysis, Atmos. Meas. Tech., 7, 1943–1955, https://doi.org/10.5194/amt-7-1943-2014, 2014.

Dusek, U., Hitzenberger, R., Kasper-Giebl, A., Kistler, M., Meijer, H. A. J., Szidat, S., Wacker, L., Holzinger, R., and Röckmann, T.: Sources and formation mechanisms of carbonaceous aerosol at a regional background site in the Netherlands: insights from a year-long radiocarbon study, Atmos. Chem. Phys., 17, 3233–3251, <u>https://doi.org/10.5194/acp-17-3233-2017</u>, 2017.

400 Elser, M., Huang, R.-J., Wolf, R., Slowik, J. G., Wang, Q., Canonaco, F., Li, G., Bozzetti, C., Daellenbach, K. R., Huang, Y., Zhang, R., Li, Z., Cao, J., Baltensperger, U., El-Haddad, I., and Prévôt, A. S. H.: New insights into PM<sub>2.5</sub> chemical composition and sources in two major cities in China during extreme haze events using aerosol mass spectrometry, Atmos. Chem. Phys., 16, 3207–3225, <u>https://doi.org/10.5194/acp-16-3207-2016</u>, 2016. Fang, W., Andersson, A., Zheng, M., Lee, M., Holmstrand, H., Kim, S.-W., Du, K., and Gustafsson, Ö.: Divergent evolution
 of carbonaceous aerosols during dispersal of East Asian haze, Sci. Rep.-UK, 7, 10422, <u>https://doi.org/10.1038/s41598-017-10766-4</u>, 2017.

Fang, W., Du, K., Andersson, A., Xing, Z., Cho, C., Kim, S. W., Deng, J., and Gustafsson, Ö.: Dual-isotope constraints on seasonally-resolved source fingerprinting of Black Carbon aerosols in sites of the four emission hotspot regions of China, J. Geophys. Res.-Atmos., 123, 11735–11747, <u>https://doi.org/10.1029/2018JD028607</u>, 2018.

410 Guo, S., Hu, M., Guo, Q., Zhang, X., Zheng, M., Zheng, J., Chang, C. C., Schauer, J. J., and Zhang, R.: Primary sources and secondary formation of organic aerosols in Beijing, China, Environ. Sci. Technol., 46, 9846–9853, 2012.

Gustafsson, Ö., Kruså, M., Zencak, Z., Sheesley, R. J., Granat, L., Engström, E., Praveen, P., Rao, P., Leck, C., and Rodhe, H.: Brown clouds over South Asia: biomass or fossil fuel combustion?, Science, 323, 495–498, 2009.

Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George,
C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, Th. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D.,
Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic aerosol: current and emerging issues,
Atmos. Chem. Phys., 9, 5155–5236, <u>https://doi.org/10.5194/acp-9-5155-2009</u>, 2009.

Heal, M. R.: The application of carbon-14 analyses to the source apportionment of atmospheric carbonaceous particulate matter: 420 a review, Anal. Bioanal. Chem., 406, 81–98, 2014.

Hu, D., Bian, Q., Lau, A. K. H., and Yu, J. Z.: Source apportioning of primary and secondary organic carbon in summer PM<sub>2.5</sub> in Hong Kong using positive matrix factorization of secondary and primary organic tracer data, J. Geophys. Res.-Atmos., 115, D16204, <u>https://doi.org/10.1029/2009jd012498</u>, 2010.

- Huang, R. J., Zhang, Y., Bozzeti, C., Ho, K. F., Cao, J. J., Han, Y., Daellenbach, K. R., Slowik, J.G., Platt, S. M., Canonaco,
  F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M., Ciarelli, G., Piazzalunga, A., Schwikowski, M., Abbaszade,
  G., SchnelleKreis, J., Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., El Haddad, I., and Prévôt, A. S. H.: High secondary aerosol contribution to particulate pollution during haze events in China, Nature, 514, 218–222, <a href="https://doi.org/10.1038/nature13774">https://doi.org/10.1038/nature13774</a>, 2014.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe,
  H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson,
  K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M.,
  Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R.,
  Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D.,
  Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel,
- 435 J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of organic aerosols in the atmosphere, Science, 326, 1525–1529, 2009.

Levin, I., Naegler, T., Kromer, B., Diehl, M., Francey, R. J., Gomez-Pelaez, A., Steele, P., Wagenbach, D., Weller, R., and Worthy, D.: Observations and modelling of the global distribution and long-term trend of atmospheric <sup>14</sup>CO<sub>2</sub>, Tellus B, 62, 26–46, 2010.

440 Lewis, C. W., Klouda, G. A., and Ellenson, W. D.: Radiocarbon measurement of the biogenic contribution to summertime PM<sub>2.5</sub> ambient aerosol in Nashville, TN, Atmos. Environ., 38, 6053–6061, 2004.

Li, C., Bosch, C., Kang, S., Andersson, A., Chen, P., Zhang, Q., Cong, Z., Chen, B., Qin, D., and Gustafsson, Ö.: Sources of black carbon to the Himalayan–Tibetan Plateau glaciers, Nat. Commun., 7, 12574, <u>https://doi.org/10.1038/ncomms12574</u>, 2016.

445 Liu, J., Li, J., Zhang, Y., Liu, D., Ding, P., Shen, C., Shen, K., He, Q., Ding, X., Wang, X., Chen, D., and Zhang, G.: Source apportionment using radiocarbon and organic tracers for PM<sub>2.5</sub> carbonaceous aerosols in Guangzhou, South China: contrasting local-and regional-scale haze events, Environ. Sci. Technol., 48, 12002–12011, 2014. Liu, J., Li, J., Liu, D., Ding, P., Shen, C., Mo, Y., Wang, X., Luo, C., Cheng, Z., Szidat, S., Zhang, Y., Chen, Y., and Zhang, G.: Source apportionment and dynamic changes of carbonaceous aerosols during the haze bloom-decay process in China based on radiocarbon and organic molecular tracers. Atmos Chem. Phys. 16, 2985–2996. https://doi.org/10.5194/gcn-16.2985-2016

450 on radiocarbon and organic molecular tracers, Atmos. Chem. Phys., 16, 2985–2996, <u>https://doi.org/10.5194/acp-16-2985-2016</u>, 2016.

Mohn, J., Szidat, S., Fellner, J., Rechberger, H., Quartier, R., Buchmann, B., and Emmenegger, L.: Determination of biogenic and fossil CO<sub>2</sub> emitted by waste incineration based on <sup>14</sup>CO<sub>2</sub> and mass balances, Bioresource Technol., 99, 6471–6479, 2008.

Ni, H., Huang, R.-J., Cao, J., Liu, W., Zhang, T., Wang, M., Meijer, H. A. J., and Dusek, U.: Source apportionment of carbonaceous aerosols in Xi'an, China: insights from a full year of measurements of radiocarbon and the stable isotope <sup>13</sup>C, Atmos. Chem. Phys., 18, 16363–16383, https://doi.org/10.5194/acp-18-16363-2018, 2018.

Ni, H., Huang, R.-J., Cao, J., Dai, W., Zhou, J., Deng, H., Aerts-Bijma, A., Meijer, H. A. J., and Dusek, U.: High contributions of fossil sources to more volatile organic aerosol, Atmos. Chem. Phys., 19, 10405–10422, https://doi.org/10.5194/acp-19-10405-2019, 2019a.

460 Ni, H., Huang, R.-J., Cao, J., Guo, J., Deng, H., and Dusek, U.: Sources and formation of carbonaceous aerosols in Xi'an, China: primary emissions and secondary formation constrained by radiocarbon, Atmos. Chem. Phys., 19, 15609–15628, https://doi.org/10.5194/acp-19-15609-2019, 2019b.

Parnell, A. C., Inger, R., Bearhop, S., and Jackson, A. L.: Source partitioning using stable isotopes: coping with too much variation, PloS ONE, 5, e9672, <u>https://doi.org/10.1371/journal.pone.0009672</u>, 2010.

465 Parnell, A. C., Phillips, D. L., Bearhop, S., Semmens, B. X., Ward, E. J., Moore, J. W., Jackson, A. L., Grey, J., Kelly, D. J., and Inger, R.: Bayesian stable isotope mixing models, Environmetrics, 24, 387–399, 2013.

Petzold, A., Ogren, J. A., Fiebig, M., Laj, P., Li, S.-M., Baltensperger, U., Holzer-Popp, T., Kinne, S., Pappalardo, G., Sugimoto, N., Wehrli, C., Wiedensohler, A., and Zhang, X.-Y.: Recommendations for reporting "black carbon" measurements, Atmos. Chem. Phys., 13, 8365–8379, <u>https://doi.org/10.5194/acp-13-8365-2013</u>, 2013.

470 Pöschl, U.: Aerosol particle analysis: challenges and progress, Anal. Bioanal. Chem., 375, 30–32, 2003.

480

Pöschl, U.: Atmospheric aerosols: composition, transformation, climate and health effects, Angew. Chem. Int. Edit., 44, 7520–7540, 2005.

Reimer, P. J., Brown, T. A., and Reimer, R. W.: Discussion: reporting and calibration of post-bomb <sup>14</sup>C data, Radiocarbon, 46, 1299–1304, 2004.

475 Santos, G. M., Southon, J. R., Griffin, S., Beaupre, S. R., and Druffel, E. R. M.: Ultra small-mass AMS <sup>14</sup>C sample preparation and analyses at KCCAMS/UCI Facility, Nucl. Instrum. Meth. B, 259, 293-302, http://dx.doi.org/10.1016/j.nimb.2007.01.172, 2007.

Szidat, S., Ruff, M., Perron, N., Wacker, L., Synal, H.-A., Hallquist, M., Shannigrahi, A. S., Yttri, K. E., Dye, C., and Simpson, D.: Fossil and non-fossil sources of organic carbon (OC) and elemental carbon (EC) in Göteborg, Sweden, Atmos. Chem. Phys., 9, 1521–1535, https://doi.org/10.5194/acp-9-1521-2009, 2009.

Szidat, S., Bench, G., Bernardoni, V., Calzolai, G., Czimczik, C. I., Derendorp, L., Dusek, U., Elder, K., Fedi, M. E., Genberg, J., Gustafsson, Ö., Kirillova, E., Kondo, M., McNichol, A. P., Perron, N., Santos, G. M., Stenström, K., Swietlicki, E., Uchida, M., Vecchi, R., Wacher, L., Zhang, Y., and Prévôt, A. S. H.: Intercomparison of <sup>14</sup>C analysis of carbonaceous aerosols: exercise 2009, Radiocarbon, 55, 1496–1509, 2013.

485 Tie, X., Huang, R.-J., Cao, J., Zhang, Q., Cheng, Y., Su, H., Chang, D., Pöschl, U., Hoffmann, T., Dusek, U., Li, G., Worsnop, D. R., and O'Dowd, C. D.: Severe pollution in China amplified by atmospheric moisture, Sci. Rep.-UK, 7, 15760, 10.1038/s41598-017-15909-1, 2017.

Turpin, B. J. and Huntzicker, J. J.: Identification of secondary organic aerosol episodes and quantitation of primary and secondary organic aerosol concentrations during SCAQS, Atmos. Environ., 29, 3527–3544, 1995.

490 van der Plicht, J., Wijma, S., Aerts, A., Pertuisot, M., and Meijer, H.: Status report: the Groningen AMS facility, Nucl. Instrum. Meth. B, 172, 58–65, 2000.

WHO: Air Quality Guidelines: Global Update 2005: Particulate Matter, Ozone, Nitrogen Dioxide and Sulfur Dioxide, World Health Organization, 2006.

Winiger, P., Andersson, A., Eckhardt, S., Stohl, A., and Gustafsson, Ö.: The sources of atmospheric black carbon at a European gateway to the Arctic, Nat. Commun., 7, 12776, <u>https://doi.org/10.1038/ncomms12776</u>, 2016.

Winiger, P., Andersson, A., Eckhardt, S., Stohl, A., Semiletov, I. P., Dudarev, O. V., Charkin, A., Shakhova, N., Klimont, Z., and Heyes, C.: Siberian Arctic black carbon sources constrained by model and observation, P. Natl. Acad. Sci. USA, 201613401, 2017.

Xi'an Municipal Bureau of Statistics and NBS Survey Office in Xi'an: Xi'an Statistical Yearbook, China Statistics Press, Beijing, China, 201709 (in Chinese).

Xu, H., Cao, J., Chow, J. C., Huang, R.-J., Shen, Z., Chen, L. A., Ho, K. F., and Watson, J. G.: Inter-annual variability of wintertime PM<sub>2.5</sub> chemical composition in Xi'an, China: evidences of changing source emissions, Sci. Total Environ., 545, 546–555, 2016.

 Zenker, K., Vonwiller, M., Szidat, S., Calzolai, G., Giannoni, M., Bernardoni, V., Jedynska, A. D., Henzing, B., Meijer, H. A.,
 and Dusek, U.: Evaluation and Inter-comparison of oxygen-based OC-EC separation methods for radiocarbon analysis of ambient aerosol particle samples, Atmosphere, 8, 226, <u>https://doi.org/10.3390/atmos8110226</u>, 2017.

Zhang, T., Cao, J.-J., Chow, J. C., Shen, Z.-X., Ho, K.-F., Ho, S. S. H., Liu, S.-X., Han, Y.-M., Watson, J. G., Wang, G.-H., and Huang, R.-J.: Characterization and seasonal variations of levoglucosan in fine particulate matter in Xi'an, China, J. Air Waste Manage., 64, 1317–1327, https://doi.org/10.1080/10962247.2014.944959, 2014.

510 Zhang, X. Y., Wang, Y. Q., Zhang, X. C., Guo, W., and Gong, S. L.: Carbonaceous aerosol composition over various regions of China during 2006, J. Geophys. Res., 113, D14111, doi:10.1029/2007JD009525, 2008.

515

Zhang, Y. L., Huang, R. J., El Haddad, I., Ho, K. F., Cao, J. J., Han, Y., Zotter, P., Bozzetti, C., Daellenbach, K. R., Canonaco, F., Slowik, J. G., Salazar, G., Schwikowski, M., Schnelle-Kreis, J., Abbaszade, G., Zimmermann, R., Baltensperger, U., Prévôt, A. S. H., and Szidat, S.: Fossil vs. non-fossil sources of fine carbonaceous aerosols in four Chinese cities during the extreme winter haze episode of 2013, Atmos. Chem. Phys., 15, 1299–1312, https://doi.org/10.5194/acp-15-1299-2015, 2015.

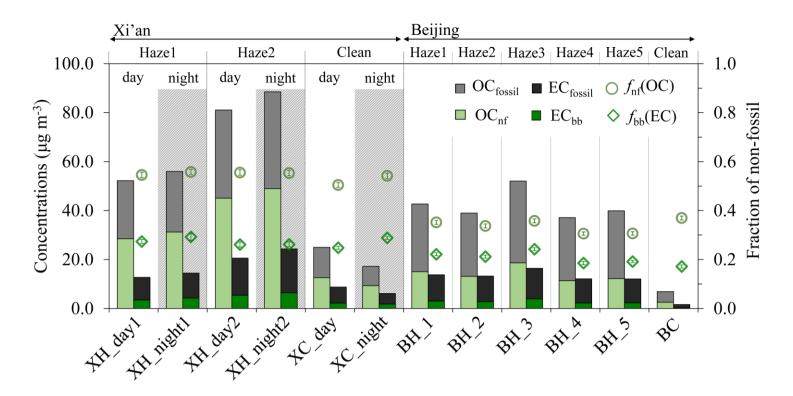
 Zheng, G. J., Duan, F. K., Su, H., Ma, Y. L., Cheng, Y., Zheng, B., Zhang, Q., Huang, T., Kimoto, T., Chang, D., Pöschl, U.,
 <u>Cheng, Y. F., and He, K. B.: Exploring the severe winter haze in Beijing: the impact of synoptic weather, regional transport</u> and heterogeneous reactions, Atmos. Chem. Phys., 15, 2969–2983, https://doi.org/10.5194/acp-15-2969-2015, 2015.

Zotter, P., Ciobanu, V. G., Zhang, Y. L., El-Haddad, I., Macchia, M., Daellenbach, K. R., Salazar, G. A., Huang, R.-J., Wacker,

520 L., Hueglin, C., Piazzalunga, A., Fermo, P., Schwikowski, M., Baltensperger, U., Szidat, S., and Prévôt, A. S. H.: Radiocarbon analysis of elemental and organic carbon in Switzerland during winter-smog episodes from 2008 to 2012 – Part 1: Source apportionment and spatial variability, Atmos. Chem. Phys., 14, 13551–13570, https://doi.org/10.5194/acp-14-13551-2014, 2014. **Table 1.** Equations for <sup>14</sup>C source apportionment of EC and OC. See Sect. 2.4 for details.  $r_{bb}$  and  $r_{fossil}$  are primary OC/EC ratio for biomass burning and fossil fuel combustion, respectively. Estimation of  $r_{bb}$  and  $r_{fossil}$  is presented in Supplement S1S2.

- L'AU	וומנ	ons

1	
$EC_{bb} = EC \times f_{bb}(EC)$	(3)
$EC_{fossil} = EC \times (1 - f_{bb}(EC)) = EC \times f_{fossil}(EC)$	(4)
$OC_{nf} = OC \times f_{nf}(OC)$	(5)
$OC_{fossil} = OC \times (1 - f_{nf}(OC)) = OC \times f_{fossil}(OC)$	(6)
$POC_{bb} = EC_{bb} \times r_{bb}$	(7)
$OC_{o,nf} = OC_{nf} - POC_{bb}$	(8)
$POC_{fossil} = EC_{fossil} \times r_{fossil}$	(9)
$SOC_{fossil} = OC_{fossil} - POC_{fossil}$	(10)
$SOC = SOC_{nf} + SOC_{fossil} \cong OC_{o,nf} + SOC_{fossil}$	(11)
$f_{\text{fossil}}(\text{SOC}) = \text{SOC}_{\text{fossil}}/\text{SOC}$	(12)



- Figure 1. Mass concentrations of OC and EC from fossil and non-fossil sources (OC<sub>fossil</sub>, OC<sub>nf</sub>, EC<sub>fossil</sub> and EC<sub>bb</sub>) as well as fraction of non-fossil carbon in OC and EC (*f*<sub>nf</sub>(OC) and *f*<sub>bb</sub>(EC), respectively) for daytime and nighttime PM<sub>2.5</sub> samples in Xi'an, and 24h-integrated PM<sub>2.5</sub> samples in Beijing during haze and clean periods <u>during the measurement periods (2 December 2016 to 10 January 2017)</u>. For each city "haze" and "clean" are used to represent high and low pollution events, and clean days at each site are defined as days with PM<sub>2.5</sub> < median concentration in the winter heating season from 15 November 2016 to 15</li>
   March 2017. Uncertainties of <sup>14</sup>C-apportioned *f*<sub>nf</sub>(OC) and *f*<sub>bb</sub>(EC) are indicated but are too small to be visible. The data are
- shown in Table S3.

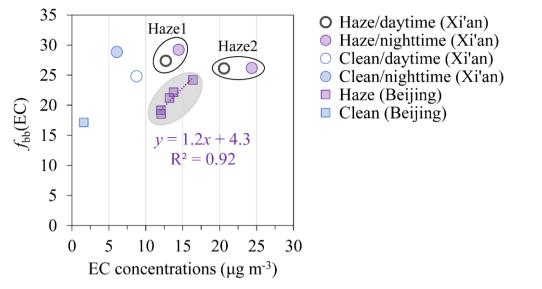
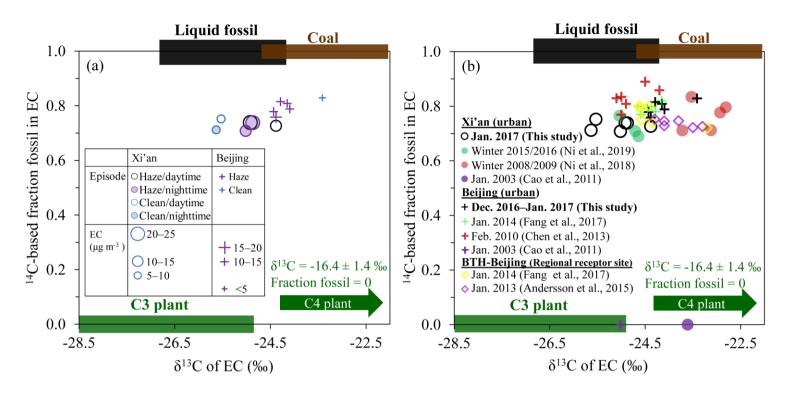
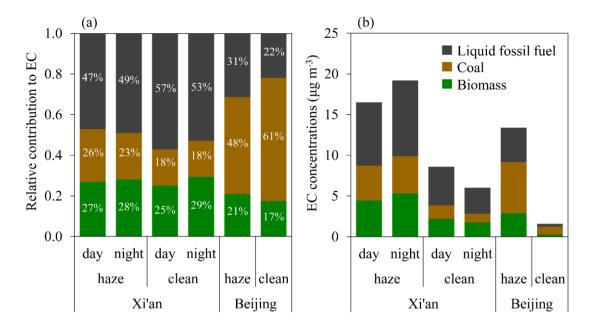


Figure 2. Scatter plot of the relative contribution of biomass burning to EC ( $f_{bb}(EC)$ ; %) against EC concentrations.

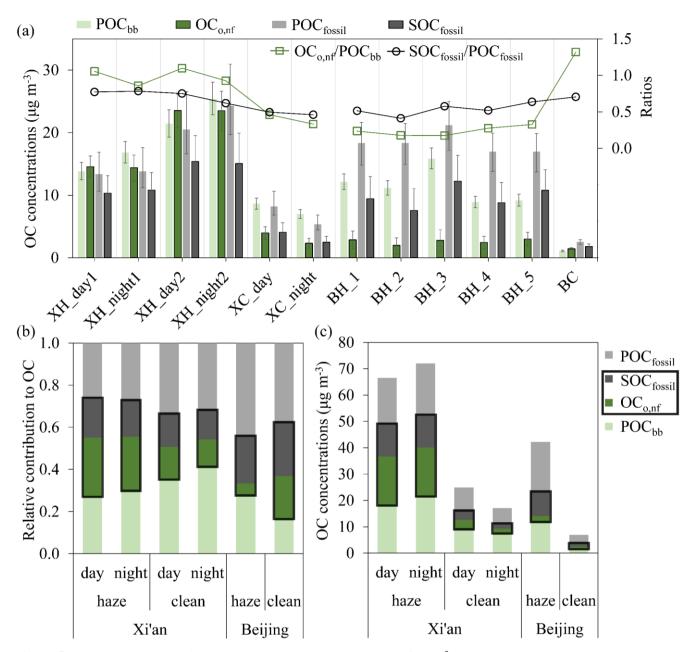


**Figure 3.** (a) <sup>14</sup>C-based fraction fossil versus  $\delta^{13}$ C for EC during haze and clean periods in Xi'an and Beijing, China. The symbol size is an indicator of EC concentrations. (b) Comparison with previous observations in Xi'an and Beijing, where BTH-Beijing is a regional receptor site of Beijing, located at 100 km southwest of Beijing. Samples from Cao et al. (2011) are placed on the *x*-axis, because no <sup>14</sup>C data were available. The expected <sup>14</sup>C and  $\delta^{13}$ C endmember ranges for emissions from C3 plant burning, liquid fossil fuel burning and coal burning are shown as green, black and brown bars, respectively. The  $\delta^{13}$ C

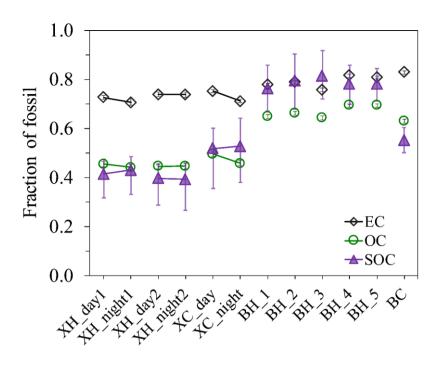
source signatures are indicated as mean  $\pm$  SD (Sect. 2.4). The  $\delta^{13}$ C signature of corn stalk burning (i.e., C4 plant; -16.4  $\pm$  1.4 ‰) is also indicated.



**Figure 4**. (a) Fractional contributions of three combustion sources to EC during haze and clean periods in Xi'an and Beijing. (b) EC concentrations ( $\mu$ g m<sup>-3</sup>) from each combustion source. The data are presented in Tables S5 and S6.



**Figure 5.** (a) Concentrations of POC<sub>bb</sub>, OC<sub>o,nf</sub>, POC<sub>fossil</sub> and SOC<sub>fossil</sub> ( $\mu$ g m<sup>-3</sup>), and the mass ratio of OC<sub>o,nf</sub>/POC<sub>bb</sub> and SOC<sub>fossil</sub>/POC<sub>fossil</sub> of each sample. The error bars indicate the interquartile range (25th–75th percentile) of the median concentrations. Averaged fraction (b) and concentration (c) of POC<sub>bb</sub>, OC<sub>o,nf</sub>, POC<sub>fossil</sub> and SOC<sub>fossil</sub> in total OC during haze and clean periods in Xi'an and Beijing, China. The data are given in Table S4.



**Figure 6.** Fraction of fossil carbon in EC, OC and SOC ( $f_{fossil}(EC)$ ,  $f_{fossil}(OC)$  and  $f_{fossil}(SOC)$ , respectively). Interquartile ranges (25th–75th percentile) of the median  $f_{fossil}(SOC)$  are shown as vertical bars in purple. Uncertainties of <sup>14</sup>C-apportioned  $f_{fossil}(EC)$  and  $f_{fossil}(OC)$  are indicated but are too small to be visible.

# Supplement of

## Measurement report: Dual-carbon isotopic characterization of carbonaceous aerosol-<u>in Beijing and Xi'an: distinctions in</u> <u>reveal different primary versus and secondary sources in</u> <u>Beijing and Xi'an during severe haze events</u>

Haiyan Ni<sup>1,2</sup>, Ru-Jin Huang<sup>1,3</sup>, Max M. Cosijn<sup>2</sup>, Lu Yang<sup>1</sup>, Jie Guo<sup>1</sup>, Junji Cao<sup>1,3</sup>, and Ulrike Dusek<sup>2</sup>

<sup>1</sup>State Key Laboratory of Loess and Quaternary Geology, Center for Excellence in Quaternary Science and Global Change, Key Laboratory of Aerosol Chemistry and Physics, Institute of Earth Environment, Chinese Academy of Sciences, Xi'an, 710061, China

<sup>2</sup>Centre for Isotope Research (CIO), Energy and Sustainability Research Institute Groningen (ESRIG), University of Groningen, Groningen, 9747 AG, the Netherlands

<sup>3</sup>Institute of Global Environmental Change, Xi'an Jiaotong University, Xi'an, 710049, China

Correspondence: Ru-Jin Huang (rujin.huang@ieecas.cn)

#### S1. Evaluation of potential interference from carbonates

PM<sub>2.5</sub> samples with the highest concentrations of calcium (associate with carbonate in most cases; Chow and Watson, 2002) in Xi'an, China were used to check the presence of carbonate carbon. Carbonate carbon in an aerosol sample can be verified if there is significant difference in TC mass before and after acidification (i.e., expose the sample to HCl vapor; NIOSH, 1999). Negligible carbonate carbon was found, because TC mass before and after acidification agrees each other within the measurement uncertainty. PM<sub>2.5</sub> samples in Beijing are less likely affected by carbonates, owing to lower dust concentrations (i.e., a major source of carbonate) than in Xi'an (Huang et al., 2014). Therefore, acidification to remove interferences from carbonates is not necessary, because carbonate carbon is very small compared to the relatively larger OC and EC amounts for both mass determination and carbon isotopic analysis.

## S12. Estimation of *r*<sub>bb</sub> and *r*<sub>fossil</sub>

 $r_{\text{bb}}$  and  $r_{\text{fossil}}$  are primary OC/EC ratio for biomass burning and fossil fuel combustion, respectively.  $r_{\text{fossil}}$  is the weighted average of (POC/EC)<sub>coal</sub> (i.e.,  $r_{\text{coal}}$ ), and (POC/EC)<sub>vehicle</sub> ( $r_{\text{vehicle}}$ ). The weight associated with  $r_{\text{coal}}$  (denoted as  $w_{\text{coal}}$ ) is the relative contribution of coal combustion to fossil EC. That is,

$$w_{\text{coal}} = \frac{\text{EC}_{\text{coal}}}{\text{EC}_{\text{fossil}}} = \frac{\text{EC}_{\text{coal}}}{\text{EC}_{\text{coal}} + \text{EC}_{\text{liq,fossil}}}$$
(S1)

where  $EC_{fossil}$  is the sum of  $EC_{coal}$  and EC from liquid fossil fuel combustion (i.e., vehicle emissions;  $EC_{liq.fossil}$ ).

Eq. (S1) can be formulated as:

$$w_{\text{coal}} = \frac{f_{\text{coal}}}{f_{\text{fossil}}} = \frac{f_{\text{coal}}}{f_{\text{coal}} + f_{\text{liq.fossil}}}$$
(S2)

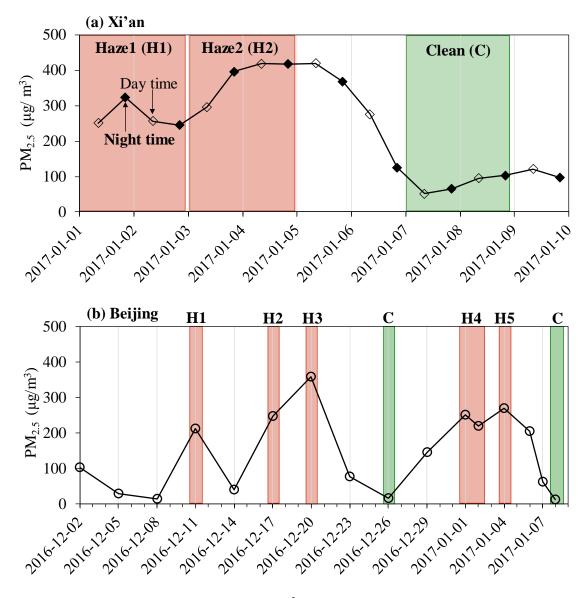
where  $f_{\text{coal}}$  and  $f_{\text{liq,fossil}}$  are the relative contribution of coal combustion and liquid fossil fuel combustion to EC, respectively. The sum of  $f_{\text{coal}}$  and  $f_{\text{liq,fossil}}$  is  $f_{\text{fossil}}$  of EC, which is well constrained by F<sup>14</sup>C of EC. The PDFs of  $f_{\text{coal}}$  and  $f_{\text{liq,fossil}}$  (e.g., Fig. S4), derived from the Bayesian calculations detailed in Sect. 2.4 of the main text, are used to calculated the PDFs of  $w_{\text{coal}}$ .

Best estimates of  $r_{bb}$  (4 ± 1; average ± SD),  $r_{coal}$  (2.38 ± 0.44), and  $r_{vehicle}$  (0.85 ± 0.16) are defined through a literature search as described in Ni et al. (2018), and their values are comparable to values used in earlier studies (Zhang et al., 2014, 2015).

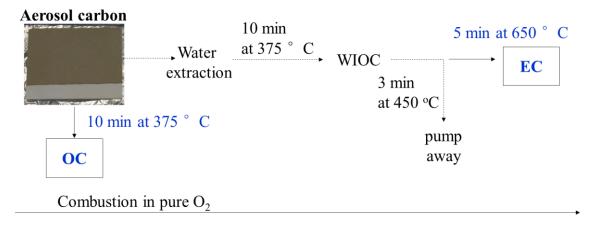
## <u>\$2</u><u>\$3</u>. Uncertainties of <sup>14</sup>C source apportionment results

To propagate uncertainties, we conducted a Monte Carlo simulation with 10000 individual calculations to propagate experimental uncertainties and uncertainties in parameters (e.g.,  $F^{14}C_{nf}$ ,  $r_{bb}$  and  $r_{fossil}$ ) following Eq. (3–12).

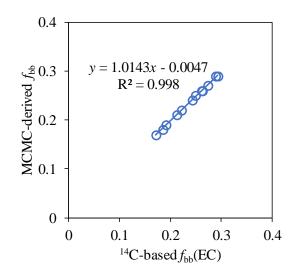
For each calculation,  $F^{14}C$  and mass of EC and OC were randomly chosen from a normal distribution symmetric around the measured values with the experimental uncertainties as standard deviation (SD). For  $F^{14}C_{nf}$ ,  $r_{bb}$ ,  $r_{coal}$  and  $r_{vehicle}$ , random values of each parameter were chosen from a triangular frequency distribution with its maximum frequency at the central value and 0 frequency at the lower limit and upper limit of each parameter. For  $w_{coal}$ , random values from the respective probability density function (PDF) of  $w_{coal}$  were used (Supplement S1S2). In this way 10000 random sets of variables were generated. For  $f_{bb}(EC)$ ,  $f_{nf}(OC)$ , EC<sub>bb</sub>, EC<sub>fossil</sub>, OC<sub>nf</sub> and OC<sub>fossil</sub>, the derived mean represents the best estimate, and the SD represents the combined uncertainties (Table S3). For POC<sub>bb</sub>, OC<sub>o,nf</sub>, POC<sub>fossil</sub>, SOC<sub>fossil</sub>, SOC and  $f_{fossil}(SOC)$ , the median value is considered as the best estimate and the interquartile ranges (25th–75th percentile) represent the combined uncertainties 4. The median values for POC<sub>bb</sub> and OC<sub>o,nf</sub> are very close to their mean values due to their symmetric PDFs (Fig. S5).



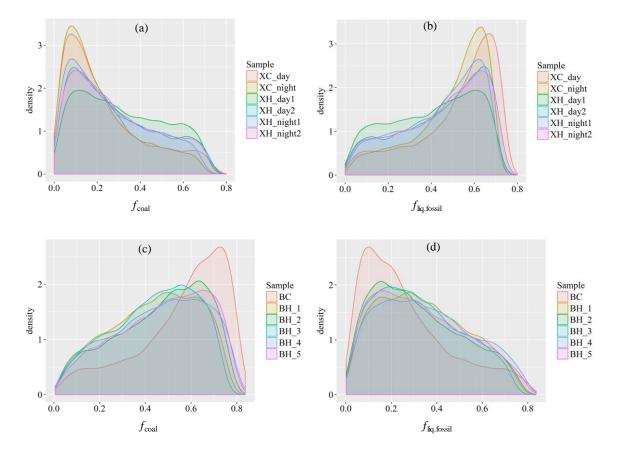
**Figure S1.**  $PM_{2.5}$  mass concentrations (µg m<sup>-3</sup>) in Xi'an and Beijing during the measurement periods. The dashed areas indicate the selected samples for <sup>14</sup>C analysis. (a) 12h averaged  $PM_{2.5}$  concentrations in Xi'an, the open symbols represent daytime samples, the filled square symbols represent nighttime samples. (b) 24h averaged  $PM_{2.5}$  concentrations in Beijing. Samples selected for <sup>14</sup>C analysis are highlighted in red (indicating haze periods) and green (clean periods). For details, see Table S1.



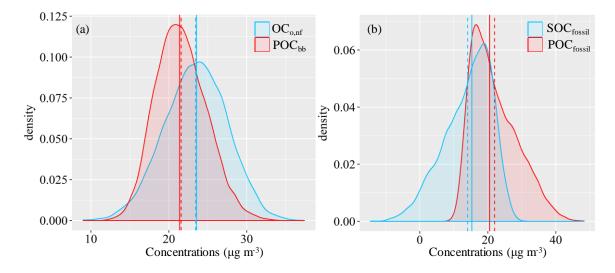
**Figure S2.** Temperature protocol of OC and EC isolation for <sup>14</sup>C measurements was implemented on aerosol combustion system (Dusek et al., 2014) in pure O<sub>2</sub>. OC is extracted by combusting the filter samples at 375 °C for 10 min. To isolate EC, OC is completely removed by 3 steps in order: water-extraction of filter samples (i.e., removal of water-soluble OC), combustion at 375 °C for 10 min (removal of water-insoluble OC; WIOC) in O<sub>2</sub>, and combustion at 450 °C for 3 min in O<sub>2</sub> (removal of the most refractory OC). Then, EC is isolated by heating the remaining carbon at 650 °C for 5 min in O<sub>2</sub>. Details can be found in Zenker et al. (2017) and Ni et al. (2018).



**Figure S3.** Comparison between the MCMC-derived fraction of biomass burning EC ( $f_{bb}$  derived from MCMC; median) and that obtained from radiocarbon data (<sup>14</sup>C-based  $f_{bb}$ (EC); mean).



**Figure S4.** Probability density functions (PDFs) of the relative source contribution of coal combustion ( $f_{coal}$ ) in Xi'an (**a**) and Beijing (**c**). PDFs of the relative source contribution of liquid fossil fuel combustion ( $f_{liq,fossil}$ ) in Xi'an (**b**) and Beijing (**d**).  $f_{coal}$  and  $f_{liq,fossil}$  are constrained by combining radiocarbon and  $\delta^{13}$ C measurements of EC, calculated using the Bayesian Markov chain Monte Carlo approach. For details, see Sect. 2.4.



**Figure S5.** (a) Example probability density functions (PDFs) of concentrations of  $OC_{o,nf}$  (light blue) and  $POC_{bb}$  (red) for sample XH\_day2. (b) PDFs of concentrations of  $SOC_{fossil}$  (light blue) and  $POC_{fossil}$  (red) for the same sample. Their concentrations are estimated from <sup>14</sup>C-apportioned OC and EC using the EC tracer method (Sect. 2.4). The mean and median are indicated by the dashed and solid vertical lines, respectively.

City	Descriptions	Selected samples for <sup>14</sup> C analysis			$\mathrm{RH}^*$	Temperature	Wind speed
		Sample name	Note	Sampling Date/Time	(%)	(°C)	(m s <sup>-1</sup> )
Xi'an	PM <sub>2.5</sub> samples were collected on the rooftop (~10 m) of a two-floor	XH_day1	Haze/daytime	2017.1.1/Daytime	69 (56~91)	4.2 (-1.5~8.3)	1.0 (0.0~2.6)
				2017.1.2/Daytime			
	building located at the Institute of Earth Environment, Chinese	XH_night1	Haze/nighttime	2017.1.1/Nighttime	88 (68~96)	0.3 (-2.6~3.5)	1.3 (0.2~2.5)
	Academy of Sciences (34.2° N, 108.9° E). This site is a typical urban background site surrounded by residential and education areas.			2017.1.2/Nighttime			
		XH_day2	Haze/daytime	2017.1.3/Daytime	78 (64~96)	4.4 (-2.6~7.2)	1.0 (0.0~2.3)
				2017.1.4/Daytime			
		XH_night2	Haze/nighttime	2017.1.3/Nighttime	93 (84~97)	0.8 (-0.6~4.1)	0.6 (0.0~1.2)
				2017.1.4/Nighttime			
		XC_day	Clean/daytime	2017.1.7/Daytime	60 (40~93)	4.4 (-4.2~8.0)	1.6 (0.4~2.6)
				2017.1.8/Daytime			
		XC_night	Clean/nighttime	2017.1.7/Nighttime	85 (75~93)	-0.9 (-3.6~1.0)	1.0 (0.0~1.8)
				2017.1.8/Nighttime			
Beijing	24 h integrated $PM_{2.5}$ samples were collected on the roof of a five-story building (~20 m) at the National Centre for Nanoscience (39.99 ° N, 116.32 ° E). The sampling site is close to the fourth ring of Beijing, and surrounded by residential, commercial and traffic areas.	BH_1	Haze	2016.12.11	59 (47~72)	0.7 (-1.7~2.1)	1.1 (0.4~1.7)
		BH_2	Haze	2016.12.17	66 (38~91)	-0.7 (-5.7~5.8)	0.8 (0.0~1.9)
		BH_3	Haze	2016.12.20	78 (55~88)	0.7 (-1.2~3.7)	1.1 (0.3~2.2)
		BH_4	Haze	2017.1.1	59 (25~89)	0.0 (-5.7~8.2)	1.4 (0.3~3.1)
				2017.1.2			
		BH_5	Haze	2017.1.4	73 (34~90)	1.1 (-3.0~7.3)	1.1 (0.0~1.7)
		BC	Clean	2016.12.26	33 (22~48)	0.5 (-4.9~8.0)	1.8 (0.3~3.3)
				2017.1.8			

Table S1. Details of sampling information and selected samples for radiocarbon measurements.

\*The meteorological data (mean; minimum-maximum) is obtained from the Meteorological Institute of Shaanxi Province, Xi'an, China.

Sample name	Note	Sampling Date/Time	$F^{14}C_{(EC)}{}^{a}$	$F^{14}C_{(OC)}{}^a$	$\delta^{13}C_{EC}^{a}$
XH_day1	Haze/daytime	2017.1.1/Daytime	$0.301\pm0.007$	$0.594 \pm 0.004$	$-24.38\pm0.02$
		2017.1.2/Daytime			
XH_night1	Haze/nighttime	2017.1.1/Nighttime $0.321 \pm 0.01$		$0.608 \pm 0.005$	$-25.02\pm0.01$
		2017.1.2/Nighttime			
XH_day2	Haze/daytime	2017.1.3/Daytime	$0.287 \pm 0.009$	$0.605\pm0.004$	$-24.92\pm0.04$
		2017.1.4/Daytime			
XH_night2	Haze/nighttime	2017.1.3/Nighttime	$0.288 \pm 0.010$	$0.603 \pm 0.004$	$-24.87\pm0.03$
		2017.1.4/Nighttime			
XC_day	Clean/daytime	2017.1.7/Daytime	$0.273 \pm 0.005$	$0.550\pm0.004$	$-25.53\pm0.02$
		2017.1.8/Daytime			
XC_night	Clean/nighttime	2017.1.7/Nighttime	$0.317 \pm 0.004$	$0.590\pm0.004$	$-25.63\pm0.03$
		2017.1.8/Nighttime			
BH_1	Haze	2016.12.11	$0.243 \pm 0.005$	$0.383 \pm 0.005$	$-24.43\pm0.03$
BH_2	Haze	2016.12.17	$0.233 \pm 0.005$	$0.368 \pm 0.005$	$-24.09 \pm 0.04$
BH_3	Haze	2016.12.20	$0.266\pm0.005$	$0.390\pm0.004$	$-24.38\pm0.02$
BH_4	Haze	2017.1.1	$0.203 \pm 0.005$	$0.333 \pm 0.004$	$-24.29\pm0.02$
		2017.1.2			
BH_5	Haze	2017.1.4	$0.211\pm0.003$	$0.334\pm0.003$	$-24.15\pm0.01$
BC	Clean	2016.12.26	$0.188 \pm 0.006$	$0.403\pm0.005$	$-23.41\pm0.01$
		2017.1.8			

**Table S2.** Fraction modern ( $F^{14}C$ ) of elemental carbon (EC), organic carbon (OC) ( $F^{14}C_{(EC)}$  and  $F^{14}C_{(OC)}$ , respectively), and stable carbon isotopic compositions ( $\delta^{13}C$ , ‰) of EC ( $\delta^{13}C_{EC}$ ).

<sup>a</sup> Values are given in average  $\pm$  measurement uncertainty.

**Table S3.** Fraction of non-fossil carbon in EC and OC ( $f_{bb}(EC)$ ,  $f_{nf}(OC)$ ), fraction of fossil carbon in EC and OC ( $f_{fossil}(EC)$ ,  $f_{fossil}(OC)$ ), concentrations ( $\mu g m^{-3}$ ) of EC and OC from non-fossil sources (EC<sub>bb</sub> and OC<sub>nf</sub>) and fossil sources (EC<sub>fossil</sub> and OC<sub>fossil</sub>) during haze and clean periods in Xi'an and Beijing during December 2016 and January 2017. Details of samples are shown in Table S1.

Sample name	$f_{bb}(EC)$	$f_{\rm fossil}({ m EC})$	$f_{\rm nf}({\rm OC})$	$f_{\rm fossil}({\rm OC})$	EC <sub>bb</sub>	EC <sub>fossil</sub>	OC <sub>nf</sub>	OC <sub>fossil</sub>
XH_day1	$0.274\pm0.008$	$0.726\pm0.008$	$0.545\pm0.011$	$0.455 \pm 0.011$	$3.48\pm0.36$	$9.24\pm0.91$	$28.49 \pm 1.53$	$23.74 \pm 1.31$
XH_night1	$0.292\pm0.014$	$0.708\pm0.014$	$0.558 \pm 0.011$	$0.442\pm0.011$	$4.23\pm0.47$	$10.24 \pm 1.05$	$31.23 \pm 1.70$	$24.77 \pm 1.40$
XH_day2	$0.261\pm0.010$	$0.739\pm0.010$	$0.555 \pm 0.011$	$0.445 \pm 0.011$	$5.37\pm0.58$	$15.22 \pm 1.55$	$45.04 \pm 2.41$	$36.08 \pm 2.02$
XH_night2	$0.262\pm0.010$	$0.738\pm0.010$	$0.553 \pm 0.011$	$0.447 \pm 0.011$	$6.37\pm0.69$	$17.96 \pm 1.82$	$48.97 \pm 2.62$	39.55 ± 2.22
XC_day	$0.248\pm0.006$	$0.752\pm0.006$	$0.505\pm0.010$	$0.495\pm0.010$	$2.17\pm0.23$	$6.56\pm0.66$	$12.61 \pm 0.68$	$12.37 \pm 0.66$
XC_night	$0.288 \pm 0.006$	$0.712\pm0.006$	$0.542 \pm 0.011$	$0.458\pm0.011$	$1.76\pm0.18$	$4.33 \pm 0.44$	9.31 ± 0.50	$7.88 \pm 0.44$
BH_1	$0.221\pm0.006$	$0.779\pm0.006$	$0.352\pm0.008$	$0.648\pm0.008$	$3.04 \pm 0.32$	$10.71 \pm 1.08$	$15.02\pm0.83$	$27.66 \pm 1.44$
BH_2	$0.212\pm0.006$	$0.788 \pm 0.006$	$0.337 \pm 0.008$	$0.663 \pm 0.008$	$2.80\pm0.29$	$10.44 \pm 1.04$	$13.13\pm0.71$	25.81 ± 1.31
BH_3	$0.242\pm0.006$	$0.758 \pm 0.006$	$0.358\pm0.008$	$0.642\pm0.008$	$3.97\pm0.41$	$12.42 \pm 1.25$	$18.62 \pm 1.01$	33.43 ± 1.72
BH_4	$0.185\pm0.006$	$0.815\pm0.006$	$0.306\pm0.007$	$0.694\pm0.007$	$2.24\pm0.23$	$9.86\pm0.97$	$11.35\pm0.62$	25.77 ± 1.29
BH_5	$0.191\pm0.004$	$0.809\pm0.004$	$0.306\pm0.007$	$0.694\pm0.007$	$2.31\pm0.24$	$9.74\pm0.97$	$12.23\pm0.67$	$27.68 \pm 1.42$
BC	$0.171\pm0.006$	$0.829 \pm 0.006$	$0.370\pm0.008$	$0.630\pm0.008$	$0.27\pm0.03$	$1.33 \pm 0.13$	$2.54\pm0.14$	$4.34\pm0.23$

**Table S4.** Concentrations ( $\mu$ g m<sup>-3</sup>) of primary OC from biomass burning (POC<sub>bb</sub>), primary OC from fossil sources, OC from non-fossil sources excluding primary biomass burning (OC<sub>o,nf</sub>), secondary OC from fossil sources (SOC<sub>fossil</sub>) and total SOC (i.e., approximately the sum of OC<sub>o,nf</sub> and SOC<sub>fossil</sub>) during haze and clean periods in Xi'an and Beijing during December 2016 and January 2017 (median and interquartile range). Details of samples are shown in Table S1.

Sample name	POC <sub>bb</sub>	POC <sub>fossil</sub>	OC <sub>o,nf</sub>	SOC <sub>fossil</sub>	SOC	$f_{\text{fossil}}(\text{SOC})$
XH_day1	13.82	13.38	14.58	10.34	24.56	0.41
	(12.50–15.26)	(10.61–16.88)	(12.83–16.25)	(6.71–13.16)	(20.53–28.25)	(0.32–0.47)
XH_night1	16.80	13.80	14.40	10.82	24.81	0.43
	(15.14–18.58)	(11.24–17.61)	(12.27–16.41)	(7.04–13.61)	(20.35–28.80)	(0.33–0.49)
XH_day2	21.43	20.50	23.54	15.41	38.29	0.40
	(19.28–23.68)	(16.60–26.58)	(20.75–26.20)	(9.35–19.52)	(31.42–44.04)	(0.29–0.45)
XH_night2	25.40	24.32	23.51	15.06	37.85	0.39
	(22.88–28.07)	(19.66–30.96)	(20.34–26.62)	(8.37–19.94)	(30.55–44.63)	(0.27–0.46)
XC_day	8.64	8.21	3.98	4.07	7.76	0.52
	(7.77–9.54)	(6.82–10.63)	(2.95–4.94)	(1.73–5.59)	(5.18–9.89)	(0.36–0.60)
XC_night	6.99	5.37	2.32	2.47	4.57	0.53
	(6.30–7.72)	4.46-6.85	1.51-3.07	1.00-3.43	2.83-6.07	0.38–0.64
BH_1	12.15	18.32	2.87	9.44	12.33	0.76
	(10.90–13.44)	(14.82–21.65)	(1.47–4.27)	(5.86–12.99)	98.13–16.41)	(0.67–0.86)
BH_2	11.17	18.36	1.98	7.56	9.59	0.80
	(10.04–12.34)	(14.89–21.44)	(0.65–3.17)	(4.26–11.04)	(5.67–13.45)	(0.68–0.91)
BH_3	15.82	21.21	2.77	12.24	15.01	0.81
	(14.21–17.56)	(17.18–25.02)	(0.95–4.50)	(8.16–16.36)	(10.19–19.75)	(0.72–0.92)
BH_4	8.88	16.95	2.45	8.79	11.30	0.78
	(8.00–9.84)	(13.83–19.94)	(1.41–3.44)	(5.70–12.06)	(7.71–14.83)	(0.70–0.86)
BH_5	9.17	16.94	3.00	10.79	13.82	0.78
	(8.25–10.17)	(13.73–19.86)	(1.96–4.07)	(7.65–14.08)	(10.26–17.53)	(0.71–0.85)
BC	1.09	2.55	1.44	1.80	3.26	0.55
	(0.98–1.21)	(2.15–2.90)	(1.30–1.59)	(1.42–2.22)	(2.79–3.76)	(0.50–0.60)

**Table S5.** Fractional contribution of different incomplete combustion sources to EC in different seasons ( $f_{bb}$ ,  $f_{iiq.fossil}$  and  $f_{coal}$ ; median and interquartile range ), and EC concentrations ( $\mu g m^{-3}$ ) from biomass burning (EC<sub>bb</sub>), coal combustion (EC<sub>coal</sub>) and liquid fossil fuel combustion (EC<sub>liq.fossil</sub>) (median and interquartile range). Details of samples are shown in Table S1.

City	Sample Name	Note	$f_{ m bb}$	$f_{ m liq.fossil}$	$f_{\rm coal}$	EC <sub>bb</sub>	EC <sub>liq.fossil</sub>	EC <sub>coal</sub>
Xi'an	XH_day1	Haze1/daytime	0.27	0.44	0.29	3.48	5.49	3.66
			(0.27–0.28)	(0.26–0.58)	(0.14–0.47)	(3.24–3.73)	(3.10–7.30)	(1.82–5.98)
	XH_night1	Haze1/nighttime	0.29	0.47	0.23	4.23	6.75	3.29
			(0.29–0.30)	(0.30–0.59)	(0.12–0.41)	(3.93–4.54)	(4.28-8.45)	(1.69–5.87)
	XH_day2	Haze2/daytime	0.29	0.47	0.23	5.38	10.04	4.94
			(0.29–0.30)	(0.30–0.59)	(0.12–0.41)	(4.99–5.78)	(6.15–12.60)	(2.44-8.94)
	XH_night2	Haze2/nighttime	0.26	0.49	0.25	6.39	11.85	5.88
			(0.26–0.27)	(0.30–0.62)	(0.12–0.44)	(5.94–6.85)	(7.40–14.93)	(2.91–10.35)
	XC_day	Clean/daytime	0.25	0.57	0.18	2.17	4.73	1.69
			(0.24–0.26)	(0.40–0.66)	(0.09–0.35)	(2.01–2.32)	(3.23–5.68)	(0.83–3.24)
	XC_night	Clean/nighttime	0.29	0.54	0.17	1.76	3.17	1.07
			(0.28–0.30)	(0.38–0.62)	(0.09–0.33)	(1.63–1.89)	(2.21–3.76)	(0.53–2.05)
Beijing	BH_1	Haze1	0.22	0.32	0.46	3.05	4.38	6.18
			(0.21–0.23)	(0.18–0.49)	(0.29–0.60)	(2.83–3.28)	(2.49–6.68)	(3.90-8.18)
	BH_2	Haze2	0.21	0.29	0.50	2.80	3.89	6.42
			(0.21–0.22)	(0.16–0.46)	(0.33–0.62)	(2.59–3.02)	(2.15–6.13)	(4.20-8.15)
	BH_3	Haze3	0.24	0.31	0.45	3.98	5.14	7.17
			(0.23–0.25)	(0.18–0.47)	(0.29–0.58)	(3.69–4.28)	(2.99–7.79)	(4.52–9.38)
	BH_4	Haze4	0.19	0.33	0.49	2.24	3.90	5.80
			(0.18–0.19)	(0.18–0.50)	(0.31–0.63)	(2.07–2.40)	(2.21–6.00)	(3.75–7.53)
	BH_5	Haze5	0.19	0.31	0.50	2.31	3.75	5.86
			(0.19–0.20)	(0.17–0.48)	(0.33–0.63)	(2.13–2.48)	(2.09–5.87)	(3.73–7.55)
	BC	Clean	0.17	0.22	0.61	0.27	0.35	0.96
			(0.17–0.18)	(0.12–0.37)	(0.46–0.71)	(0.25–0.30)	(0.18–0.59)	(0.72–1.14)

			Biomass burning	Coal combustion	Liquid fossil fuel combustion (i.e., vehicle emissions)
Xi'an	haze	day	0.27	0.26	0.47
			(0.26–0.27)	(0.13–0.44)	(0.29–0.60)
Xi'an	haze	night	0.28	0.23	0.49
			(0.27–0.28)	(0.12–0.41)	(0.31–0.61)
Xi'an	clean	day	0.25	0.18	0.57
			(0.24–0.26)	(0.09–0.33)	(0.42–0.66)
Xi'an	clean	night	0.29	0.18	0.53
			(0.28–0.30)	(0.09–0.34)	(0.37–0.62)
Beijing	haze		0.21	0.48	0.31
			(0.20–0.22)	(0.31–0.61)	(0.18–0.48)
Beijing	clean		0.17	0.61	0.22
			(0.17–0.18)	(0.45–0.71)	(0.12–0.38)

**Table S6.** Fractional contribution of different incomplete combustion sources to EC (median and interquartile range).

## References

Chow, J. C. and Watson, J. G.: PM<sub>2.5</sub> carbonate concentrations at regionally representative Interagency Monitoring of Protected Visual Environments sites, *J. Geophys. Res.*, 107(D21), 8344, doi:10.1029/2001JD000574, 2002.

Dusek, U., Monaco, M., Prokopiou, M., Gongriep, F., Hitzenberger, R., Meijer, H. A. J., and Röckmann, T.: Evaluation of a two-step thermal method for separating organic and elemental carbon for radiocarbon analysis, Atmos. Meas. Tech., 7, 1943–1955, <u>https://doi.org/10.5194/amt-7-1943-2014</u>, 2014.

Huang, R. J., Zhang, Y., Bozzeti, C., Ho, K. F., Cao, J. J., Han, Y., Daellenbach, K. R., Slowik, J.G., Platt, S. M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M., Ciarelli, G., Piazzalunga, A., Schwikowski, M., Abbaszade, G., SchnelleKreis, J., Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., El Haddad, I., and Prévôt, A. S. H.: High secondary aerosol contribution to particulate pollution during haze events in China, Nature, 514, 218–222, https://doi.org/10.1038/nature13774, 2014.

Ni, H., Huang, R.-J., Cao, J., Liu, W., Zhang, T., Wang, M., Meijer, H. A. J., and Dusek, U.: Source apportionment of carbonaceous aerosols in Xi'an, China: insights from a full year of measurements of radiocarbon and the stable isotope <sup>13</sup>C, Atmos. Chem. Phys., 18, 16363–16383, https://doi.org/10.5194/acp-18-16363-2018, 2018.

NIOSH: Method 5040 Issue 3 (Interim): Elemental carbon (diesel exhaust), in: NIOSH Manual of Analytical Methods; National Institute of Occupational Safety and Health: Cincinnati, OH, 1999.

Zenker, K., Vonwiller, M., Szidat, S., Calzolai, G., Giannoni, M., Bernardoni, V., Jedynska, A. D., Henzing, B., Meijer, H. A., and Dusek, U.: Evaluation and Inter-comparison of oxygen-based OC-EC separation methods for radiocarbon analysis of ambient aerosol particle samples, Atmosphere, 8, 226, <u>https://doi.org/10.3390/atmos8110226</u>, 2017.

Zhang, Y. L., Li, J., Zhang, G., Zotter, P., Huang, R.-J., Tang, J.-H., Wacker, L., Prévôt, A. S. H., and Szidat, S.: Radiocarbon-based source apportionment of carbonaceous aerosols at a regional background site on Hainan Island, South China, Environ. Sci. Technol., 48, 2651–2659, <u>https://doi.org/10.1021/es4050852</u>, 2014.

Zhang, Y. L., Huang, R. J., El Haddad, I., Ho, K. F., Cao, J. J., Han, Y., Zotter, P., Bozzetti, C., Daellenbach, K. R., Canonaco, F., Slowik, J. G., Salazar, G., Schwikowski, M., Schnelle-Kreis, J., Abbaszade, G., Zimmermann, R., Baltensperger, U., Prévôt, A. S. H., and Szidat, S.: Fossil vs. non-fossil sources of fine carbonaceous aerosols in four Chinese cities during the extreme winter haze episode of 2013, Atmos. Chem. Phys., 15, 1299–1312, <u>https://doi.org/10.5194/acp-15-1299-2015</u>, 2015.