



1 **Rapid mass growth and enhanced light extinction of atmospheric aerosols during the heating**
2 **season haze episodes in Beijing revealed by aerosol-chemistry-radiation-boundary layer**
3 **interaction**

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26



27 **Abstract**

28

29 Despite the numerous studies investigating haze formation mechanism in China, it is still puzzling
30 that intensive haze episodes could form within hours directly following relatively clean periods. Haze
31 has been suggested to be initiated by the variation of meteorological parameters and then to be
32 substantially enhanced by aerosol-radiation-boundary layer feedback. However, knowledge on the
33 detailed chemical processes and the driving factors for extensive aerosol mass accumulation during
34 the feedback is still scarce. Here, the dependency of the aerosol number size distribution, mass
35 concentration and chemical composition on the daytime mixing layer height (MLH) in urban Beijing
36 is investigated. The size distribution and chemical composition-resolved dry aerosol light extinction
37 is also explored. The results indicate that the aerosol mass concentration and fraction of nitrate
38 increased dramatically when the MLH decreased from high to low conditions, corresponding to
39 relatively clean and polluted conditions, respectively. Particles having their dry diameters in the size
40 of ~400-700 nm, and especially particle-phase ammonium nitrate and liquid water, contributed
41 greatly to visibility degradation during the winter haze periods. The dependency of aerosol
42 composition on the MLH revealed that ammonium nitrate and aerosol water content increased the
43 most during low MLH conditions, which may have further triggered enhanced formation of sulphate
44 and organic aerosol via heterogeneous reactions. As a result, more sulphate, nitrate and water soluble
45 organics were formed, leading to an enhanced water uptake ability and increased light extinction by
46 the aerosols. The results of this study contribute towards a more detailed understanding of the aerosol-
47 chemistry-radiation-boundary layer feedback that is likely to be responsible for explosive aerosol
48 mass growth events in urban Beijing.

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55 1. Introduction

56 Despite the recent reduction of air pollutants and their precursors in China between 2013 and 2017,
57 the current emission and air pollution levels are still substantially high (Wang et al., 2020b; Zheng et
58 al., 2018). Such high emissions, combined with specific meteorological conditions, frequently lead
59 to severe haze episodes (An et al., 2019; Wang et al., 2019). Particulate matter, a major air pollutant,
60 has considerable effects on climate, human health and visibility degradation (Che et al., 2007;
61 Lelieveld et al., 2015; Spracklen et al., 2008; Wang et al., 2015).

62

63 During winter haze episodes, a rapid growth of the aerosol mass concentration has commonly been
64 observed, and this phenomenon seems to be directly affected by meteorological factors (Li et al.,
65 2018b; Liu et al., 2018, 2019b; Wang et al., 2018a, 2014a). The meteorological conditions and
66 increased aerosol concentrations are proposed to be interlinked by a feedback loop, called the aerosol-
67 chemistry-boundary layer feedback, in which aerosol particles reduce both solar radiation reaching
68 the surface and turbulent kinetic energy of the near-surface air (Ding et al., 2016; Petäjä et al., 2016;
69 Wang et al., 2020d). The increased stability of the boundary layer leads to enhanced air pollution in
70 the mixed layer, which further suppresses the development of boundary layer. As a consequence,
71 concentrations of primary aerosol particles, water vapor and relative humidity increase, creating more
72 favourable conditions for homogeneous and heterogeneous on aerosol surfaces or inside them (Cheng
73 et al., 2016a; Wang et al., 2016; Wu et al., 2018). Such reactions cause rapid formation of secondary
74 aerosol matter and enhanced light extinction during severe winter haze episodes. However, more
75 detailed information on the aerosol and reactive gas chemistry during the aerosol-chemistry-boundary
76 layer feedback and related rapid aerosol mass growth events is still needed (Liu et al., 2019). For
77 instance, it is still unclear which chemical reactions and which compounds in the particulate matter
78 play key roles during such rapid mass growth events.

79

80 The particle number size distribution and chemical composition are considered to be the most
81 important variables influencing the light extinction by aerosol particles. In the atmosphere, the highest
82 contribution to aerosol light extinction comes from organic compounds, nitrate and sulphate in



83 particles with diameters of a few hundred nm. This is due to the dominant mass fractions of the
84 aforementioned compounds in aerosols that correspond to the peak intensity of solar radiation at
85 wavelengths around 550 nm (Jimenez et al., 2009; Swietlicki et al., 2008). In addition, light scattering
86 which contributes the most to the light extinction by atmospheric aerosols, can be substantially
87 enhanced by the presence of liquid water in the aerosol (Chen et al., 2014; Liu et al., 2019a; Pan et
88 al., 2009; Wang et al., 2020). Hence, quantifying the response of light extinction to different chemical
89 compounds would be helpful in evaluating the feedbacks associated with secondary aerosol
90 production.

91

92 In this study, we focus on the physical and chemical properties of aerosols in Beijing during the winter
93 heating season from October 2018 to February 2019 using state-of-the-art instrumentation. The
94 variation of aerosol chemical composition and the associated light extinction coefficient as a function
95 of the varying mixing layer height are discussed. Our aim is to identify the key chemical components
96 which contribute to the aerosol-chemistry-radiation-boundary layer feedback loop in Beijing.

97

98 2. Methodology

99 2.1. Measurement location and instrumentations

100 Measurements were conducted between 1 October 2018 and 28 February 2019 at the roof top of the
101 university building at the west campus of Beijing University of Chemical Technology (39.95°N,
102 116.31°E). This station is located about 150 m away from the nearest road (Zizhuyuan road) and 500
103 m away from the West Third Ring Road, and it is surrounded by commercial properties and residential
104 dwellings representative of an urban environment. More details on the location can be found in (Liu
105 et al., 2020; Zhou et al., 2020).

106

107 The meteorological data for this work include basic meteorological variables (relative humidity (RH),
108 temperature, wind speed, wind direction, and visibility) and mixing layer height (MLH) measured



109 using a weather station (Vaisala Inc., Finland) and a Ceilometer CL51 (Vaisala Inc., Finland),
110 respectively. The MLH is defined as the height above the surface, through which relatively vigorous
111 vertical mixing occurs (Holzworth, 1972), and its value is highly related to the vertical temperature
112 structure and, so some extent, to a mechanically-induced turbulence (Baxter, 1991). Here, we
113 followed the method introduced earlier by Mönkel et al. (2007) and Eresmaa et al. (2012) in
114 determining the MLH.

115

116 The number concentration of clusters or small aerosol particles in the size range from 1.3-2.5 nm and
117 the number size distributions of aerosol particles from 6 nm to 840 nm were measured by a Particle
118 Sizer Magnifier (PSM) and a Differential Mobility Particle Sizer (DMPS), respectively (Aalto et al.,
119 2001; Vanhanen et al., 2011). The mass concentration of fine particulate matter (PM_{2.5}) was measured
120 using a Tapered Element Oscillating Microbalance Dichotomous Ambient Particulate Monitor
121 (TEOM 1405-DF, Thermo Fisher Scientific Inc, USA) with a total flow rate of 16.67 L/min (Wang
122 et al., 2014).

123

124 A time-of-flight aerosol chemical speciation monitor (ToF-ACSM, Aerodyne Research Inc.) was used
125 to measure the concentrations of non-refractory (NR) components, including sulfate, nitrate,
126 ammonium, chloride and organics of PM_{2.5} (Fröhlich et al., 2013). The inlet flow was set at 1.4 cm³/s.
127 The particle beam passed through the chamber and reached the heated porous tungsten surface
128 (T≈600°C). There, the non-refractory PM_{2.5} constituents were vaporized and then ionized by electrons
129 (E_{kin}=70eV, emitted by a tungsten filament). The ions were measured by a detector and the data was
130 analyzed using Tofware ver. 2.5.13 within IgorPro ver. 6.3.7.2 (WaveMetrics). The relative ionization
131 efficiencies (RIE) for sulfate, nitrate, ammonium, chloride and organics applied were 0.86, 1.05, 4.0,
132 1.5 and 1.4, respectively. In addition to the RIE corrections, CO₂⁺/NO₃ artifact correction and
133 collection efficiency (CE) correction (Pieber et al., 2016) were also applied to the data. Mass
134 concentrations of ammonium nitrate, ammonium sulfate and ammonium chloride were determined
135 according to the method introduced by Gysel et al. (2007). The aerosol liquid water content (AWC)
136 was calculated by thermodynamic equilibrium model ISORROPIA II using ToF-ACSM data



137 (Fountoukis and Nenes, 2007).

138

139 Highly-oxygenated organic molecules (HOMs) were measured by a chemical ionization long time-
140 of-flight mass spectrometer equipped with a nitrate chemical ionization source (LToF-CIMS,
141 Aerodyne Research, Inc. USA) (Jokinen et al., 2012) similar to gas-phase sulfuric acid. Organic
142 carbon (OC) and element carbon (EC) concentrations were measured semi-continuously with a 1-
143 hour time resolution using an OC/EC Analyzer (Model-4, Sunset Lab. Inc.).

144

145 The air mass history was studied by calculating particle retroplumes using a Lagrangian particle
146 dispersion model FLEXPART (FLEXible PARTicle dispersion model) ver. 9.02 (Stohl et al., 2005).
147 The ECMWF (European Centre for Medium-Range Weather Forecast) operational forecast (with 0.15°
148 horizontal and 1 h temporal resolution) was used as the meteorological input into the model. During
149 the measurement period, a new release of 50 000 test particles, distributed evenly between 0 and 100
150 m above the measurement site, occurred every 1 hour. The released particles were traced backwards
151 in time for 72 h, unless they exceeded the model boundary (20–60°N, 95–135°E).

152

153 **2.2. Aerosol light extinction calculation**

154 The aerosol light extinction coefficient was calculated with the Mie-Model, which uses particle
155 number size distribution, mass concentrations of different aerosol compounds and their refractive
156 index as inputs (Seinfeld and Pandis, 2006). We introduced a series of assumptions into the Mie-
157 Model, including 1) “internal mixture” which considers each chemical component in a particle as
158 homogeneously mixed with each other; 2) all particles are spherical; and 3) particles of different sizes
159 have the same chemical composition.

160

161 The practical method introduced under those assumptions in previous studies were found to be
162 capable of estimating a variation trend of optical property of PM_{0.5–20} with a relatively good accuracy
163 (Lin et al., 2013).



164

165 Table 1. Summary of the parameters for calculating the average optical refractive index.

166

Specie	$\rho_i(\text{g cm}^{-3})$	n_i	k_i
$(\text{NH}_4)_2\text{SO}_4$	1.760	1.530	0.000
NH_4NO_3	1.725	1.554	0.000
NH_4Cl	1.527	1.639	0.000
Organics	1.400	1.550	0.001
EC	1.500	1.800	0.540

167

168 The average optical refractive index (AORI) of an internally-mixed particle can be calculated from
 169 the optical refractive indices (ORI) of each chemical component by following a mixing rule of
 170 volume-averaged chemical components as $\text{AORI} = n_{\text{eff}} + k_{\text{eff}} \times i$, where the real part (n_{eff}) and
 171 imaginary part (k_{eff}) are given by:

$$n_{\text{eff}} = \left(\sum_i n_i \cdot m_i / \rho_i \right) / \left(\sum_i m_i / \rho_i \right) \quad (1)$$

$$k_{\text{eff}} = \left(\sum_i k_i \cdot m_i / \rho_i \right) / \left(\sum_i m_i / \rho_i \right) \quad (2)$$

172 Here m_i and ρ_i are the mass concentration and density of the component i in particles, respectively,
 173 and n_i and k_i are the real and imaginary parts of ORI of this component, respectively. The
 174 parameters for calculating the AORI are summarised in Table 1. The values of n_i and k_i in Table 1
 175 are referenced to the light wavelength of 550 nm.

176

177 $Q_{sp,j}$ represents light scattering efficiency of a single particle with diameter D_j , while $Q_{ep,j}$
 178 represents light absorption efficiency. Theoretically, $Q_{sp,j}$ and $Q_{ep,j}$ are both the function of D_j and
 179 the AORI_j (the AORI of the particle with diameter D_j) at a given light wavelength λ , for which the
 180 complicated calculations were referenced to a previous publication. Regarding the limitations of
 181 measurement techniques, the AORI_j was assumed to be equal to the $\text{AORI}_{\text{PM}_{2.5}}$, which was



182 determined based on chemical composition of PM_{2.5}. It is possible to derive expressions for the cross
 183 sections of a spherical particle exactly. The formulas for $Q_{sp,j}$ and $Q_{ep,j}$ are:

184

$$Q_{sp,j}(D_j, \lambda, AORI_j) = \frac{2}{\alpha^2} \sum_{k=1}^{\infty} (2k+1) \cdot [|a_k|^2 + |b_k|^2] \quad (3)$$

$$Q_{ep,j}(D_j, \lambda, AORI_j) = \frac{2}{\alpha^2} \sum_{k=1}^{\infty} (2k+1) \cdot \text{Re}[a_k + b_k] \quad (4)$$

185

186 where

187

$$a_k = \frac{\alpha \psi'_k(y) \psi_k(\alpha) - y \psi'_k(\alpha) \psi_k(y)}{\alpha \psi'_k(y) \xi_k(\alpha) - y \xi'_k(\alpha) \psi_k(y)}$$

189

$$b_k = \frac{y \psi'_k(y) \psi_k(\alpha) - \alpha \psi'_k(\alpha) \psi_k(y)}{y \psi'_k(y) \xi_k(\alpha) - \alpha \xi'_k(\alpha) \psi_k(y)}$$

191

192 with $y = \alpha m$.

193

$$m = n_{eff} + i \cdot k_{eff}$$

195

$$\alpha = \frac{\pi D_j}{\lambda}$$

197

198 with $\lambda = 550$ nm.

199

200 where complex number m stands for $AORI_j$, while α is the size of the particle, usually expressed as

201 a dimensionless size parameter. The functions $\psi_k(z)$ and $\xi_k(z)$ are the Riccati–Bessel functions:

$$\psi_k(z) = \left(\frac{\pi z}{2}\right)^{1/2} J_{k+1/2}(z) \quad (5)$$

$$\xi_k(z) = \left(\frac{\pi z}{2}\right)^{1/2} [J_{k+1/2}(z) + i(-1)^k J_{-k-1/2}(z)] \quad (6)$$

202



where $J_{k+1/2}$ and $J_{-k-1/2}$ are the Bessel functions of the first kind and their footnotes indicate the order of Bessel functions. The Mie theory can serve as the basis of a computational procedure to calculate the scattering and absorption of light by any sphere as a function of wavelength.

According to the Mie-Model, b_{sp} (light scattering coefficient) and b_{ep} (light extinction coefficient) can be quantified with Eqs. (5) and (6), respectively. b_{ap} (light absorption coefficient) is the difference between b_{ep} and b_{sp} , which equals zero, when k_i equals zero or very small. Optical properties including b_{ep} , b_{sp} and b_{ap} to be discussed later are all referenced to light wavelength of 550 nm.

$$b_{sp} = \sum_j b_{sp,j} = \sum_j \frac{\pi D_j^2}{4} \cdot Q_{sp,j}(D_j, \lambda, AORI_j) \cdot N_j \quad (7)$$

$$b_{ep} = \sum_j b_{ep,j} = \sum_j \frac{\pi D_j^2}{4} \cdot Q_{ep,j}(D_j, \lambda, AORI_j) \cdot N_j \quad (8)$$

In Eqs. (7) and (8), D_j stands for the median Stokes diameter in the j -th particle size range and N_j is the number concentration of particles with diameter, D_j .

3. Results and discussion

3.1. Typical case of rapid aerosol mass growth episodes affected by aerosol-chemistry-boundary layer interactions

An example of rapid aerosol mass growth in urban wintertime Beijing is illustrated in Figure 1, where the haze accumulation was associated with a rapid $PM_{2.5}$ mass concentration increase from few $\mu g/m^3$ to more than 100 $\mu g/m^3$ in less than 7 hours. A haze episode started on afternoon 20 February 2019 under stagnant meteorological conditions with low wind speeds and elevated ambient relative humidity (Figure S1). The polluted periods during this case occurred under southerly wind transport



225 conditions, whereas clean air masses originated from the north-westerly regions (as shown in Figure
226 S2, S3). These are typical features for a haze evolution process in Beijing (Wang et al., 2020b). During
227 the haze periods marked by the shaded areas in Figure 1, an obvious increase of chemical mass
228 concentration was observed by the ToF-ACSM, characterised by high concentrations of secondary
229 aerosol components (nitrate, organics and sulphate) and typically a shallow boundary layer. The mass
230 concentrations of organics, sulphate and nitrate increased dramatically with a decreasing MLH,
231 accounting for 88.5% of NR-PM_{2.5} during the rapid aerosol mass growth period. The aerosol mass
232 growth was the fastest for nitrate. The mass concentrations of organic and elemental carbon followed
233 that of NR-PM_{2.5}.

234

235 The MLH reached its maximum at around 14:00 in the afternoon of 20 February, after which the
236 development of the mixing layer was suppressed and MLH decreased with the arrival of pollution
237 (Figure 1a). Previous studies have shown that the aerosol-radiation-boundary layer feedback
238 contributes to a rapid enhancement of air pollution (Petäjä et al., 2016; Wang et al., 2020d). High
239 concentrations of aerosol particles obscure downward radiation, as a result of which the surface
240 temperature and sensitive heat flux decrease and the development of mixing layer height is suppressed.
241 Recent studies have gradually realized that the facilitation of various chemical processes play a non-
242 negligible role in the aerosol-radiation-boundary layer feedback (Liu.Q et al., 2018; Liu. Z et al., 2019).
243 Therefore, it is important to identify and quantify the role of different specific chemical species and
244 particle size ranges in reducing atmospheric radiation and extinction.

245

246 Figure 2 shows the contributions of size and chemical composition-resolved dry aerosol to light
247 extinction during the investigated period. As the pollution intensified and MLH decreased (Fig 1c),
248 the light extinction of atmospheric aerosols increased significantly. Assuming that particles of
249 different sizes have the same chemical composition as PM_{2.5} (organics, NH₄NO₃, EC, (NH₄)₂SO₄,
250 NH₄Cl), the light extinction of particles in the size range of 300-700 nm increased significantly from
251 the relative clean period to the polluted period (namely from 12:00 to 16:00). During relatively clean
252 conditions, the contributions of organics, NH₄NO₃, EC, (NH₄)₂SO₄ and NH₄Cl to the total aerosol



light extinction were 42%, 23%, 18%, 11% and 7%, respectively. The contribution of NH_4NO_3 to aerosol light extinction reached 40% during the heavily polluted period. The increased light extinction by aerosols reduced solar radiation reaching the surface, so that the development of the boundary layer was suppressed.

3.2. Connection between the aerosol chemical composition, light extinction, size distribution and MLH during the heating season

To better characterize the effect of the chemical composition of dry aerosols and the PNSD (particle number size distribution) light extinction under different MLH conditions, the daytime (8:00 – 16:00 LT) measurement data from October 2018 to February 2019 were selected for further analysis. As shown by Figure 3 and consistent with other observations in Beijing (Tang et al., 2016; Wang et al., 2020c), there was a general tendency for the $\text{PM}_{2.5}$ mass concentration to increase with a decreasing MLH. Organic compounds and nitrate were the most abundant fractions of the daytime aerosol mass composition, contributing together approximately 70% to total NR- $\text{PM}_{2.5}$ mass concentration. With a decreasing MLH, the fraction of nitrate mass in NR- $\text{PM}_{2.5}$ slightly increased while that of organics decreased. This feature makes the aerosol more hygroscopic under low MLH conditions typical for heavily polluted periods. The increased nitrate fraction in the aerosol could also enhance the formation of other secondary aerosol components (Xue et al., 2019). Note that some fraction of aerosol nitrate could consist of organic nitrate originating from reaction of peroxy radical with nitric oxide; however, it is difficult to distinguish organic nitrate from inorganic nitrate at the moment due to instrumental limitations (Fröhlich et al., 2013).

Figure 4 depicts the calculated daytime light extinction of the dry aerosol as a function of the MLH, separated by different size ranges and chemical components. We may see that in general, particles with dry diameters in the range of 300-700 nm explained more than 80% of the total aerosol light extinction (Figure 4b). Similar to their share in NR- $\text{PM}_{2.5}$, the fraction of light extinction by



280 ammonium nitrate increased and that of organics decreased during the lowest MLH conditions
281 corresponding to the heavy pollution periods (Figure 4d). There are also apparent differences in the
282 relative contribution of different particle size ranges to light extinction in different MLH conditions:
283 with a decreasing MLH, the contribution of particles with dry dimeters larger than about 400-500 nm
284 clearly increased while that of sub-300 nm particles notably decreased. This indicates that the
285 enhanced light extinction by the dry aerosol at low MLH conditions was not only due the more
286 abundant aerosol mass concentration, but also due to the growth of individual particles to optically
287 more active sizes.

288

289 At relative humidity larger than about 70%, aerosol liquid water gives a significant contribution to
290 the aerosol mass concentration and often a dominant contribution to the aerosol light extinction (Titos
291 et al., 2016). This has important implications for the aerosol-chemistry-radiation-boundary layer
292 feedback, when considering our findings listed above and further noting that heavy pollution periods
293 are often accompanied by high values of RH in Beijing (Zhong et al., 2018). First, compared to clean
294 or moderately-polluted conditions, the enhancement in the aerosol light extinction under polluted is
295 probably much larger than that illustrated in Figure 4. Second, the high aerosol water content under
296 polluted conditions promotes many kinds of chemical reactions taking place on the surface or inside
297 aerosol particles.

298

299 3.3. Aerosol-chemistry-radiation-boundary layer interaction

300

301 In order to further investigate the interaction between MLH and chemical compounds (either observed
302 or calculated), we divided the observed $PM_{2.5}$ concentrations into highly polluted and less polluted
303 conditions using a threshold value of $75 \mu g m^{-3}$ for $PM_{2.5}$. The organics, nitrate, ammonium, sulfate,
304 chloride, HOM, aerosol water content (AWC) and $PM_{2.5}$ as a function of the mixing layer height
305 during both highly polluted and less polluted conditions are shown in Figure 5. The fitted relationships
306 connecting the concentrations of different chemical compounds to the reduction of MLH under highly



307 and less polluted conditions allowed us to estimate the net mass concentration increase of each
308 compound due to secondary formation and aerosol-chemical-boundary layer feedback under highly
309 polluted conditions (shaded areas in Figure 5). It is worth noting that AWC, nitrate and sulfate
310 increased the most as the MLH decreased, as represented by the large shaded areas in Figs. 5 (h), (b)
311 and (c). The day-time nitrate in aerosol is formed predominately via the reaction of nitric acid and
312 ammonium, while nitric acid is produced from gas phase reaction of nitrogen dioxide and hydroxy
313 radical (Seinfeld and Pandis, 2006). High concentrations of daytime nitrate aerosols indicate efficient
314 production of gas phase nitric acid, its partitioning into liquid aerosol and its fast neutralization by
315 abundant ammonia (Li et al., 2018a; Pan et al., 2016; Wang et al., 2020). A recent study shows that
316 condensation of nitric acid and ammonia could promote fast growth of newly formed particle in urban
317 environment condition (Wang et al., 2020d). Another possibility is that ammonium nitrate is formed
318 rapidly on particle surfaces due to the hydrolysis of dinitrogen pentoxide (N_2O_5) during daytime, as
319 the AWC increased significantly (Wang et al., 2014; Wang et al., 2020). However, a quantitative
320 distinction between the two formation pathways for nitrate formation is not possible in this study. The
321 dramatic increase of nitrate aerosol could also promote the formation of sulfate by heterogeneous
322 reactions (Cheng et al., 2016b; Wang et al., 2016). The concentration of HOMs showed a slight
323 increase as the MLH decreased, which suggests that also the formation of HOMs is enhanced with an
324 increased level of air pollution. This phenomenon should be further investigated as HOMs can
325 substantially contribute to the secondary organic aerosol formation.

326
327 Figure 6 displays the dry aerosol light extinction by different chemical compounds in the same way
328 as Fig. 5 did for aerosol mass concentrations. The aerosol light extinction is directly related to the
329 reduction of solar radiation reaching the surface, assuming that aerosol chemical components are
330 vertically nearly homogeneously distributed. The light extinction from ammonium nitrate,
331 ammonium sulfate and organics showed significantly increased contributions under highly polluted
332 conditions (low MLH) as compared with less polluted conditions. To the contrary, no such
333 enhancement was observed for ammonium chloride or element carbon (Figs. 6 (d) and (e)). In case
334 of EC this is an expected result, as it originates solely from primary sources. The formation of particle



phase chloride have secondary sources from chlorine atom-initiated oxidation of volatile organic compounds, so that the resulting oxidation products could contribute to the observed chloride (Wang and Ruiz, 2017; Wang et al., 2019a).

To better illustrate the combined effects of secondary aerosol formation and associated feedback on the daytime mass concentrations and light extinction due to different chemical components, we scaled these quantities by either the total $PM_{2.5}$ mass concentration or EC concentration and plotted them as a function of MLH (Fig. 7). The latter scaling minimizes the boundary layer accumulation effect on our analysis, as EC originates from primary emission sources (Cao et al., 2006). As shown in Fig. 7a, organics with their mass fraction of 61% were the most abundant component in $PM_{2.5}$ under high MLH conditions, followed by nitrate and ammonium with their mass fractions of 22% and 13%, respectively. The aerosol was estimated to be rather dry under high MLH conditions ($AWC/PM_{2.5} = 0.03$). However, with the decreasing MLH, the fraction of nitrate and the AWC to $PM_{2.5}$ ratio increased up to 45% and 0.2, respectively. This clearly indicates rapid nitrate formation and dramatic increase of the aerosol water uptake from less polluted conditions to intensive haze pollution. Compared with EC (Fig. 7c), the concentrations of organic compounds, nitrate, sulfate and ammonium increased by factors of 1.5, 6.3, 4.8 and 4.9 respectively, from the highest to the lowest MLH conditions. Thus, although organics remained as the second most abundant aerosol component after nitrate under haze conditions, secondary formation and associated feedback from less to highly polluted conditions were clearly stronger for both sulfate and ammonium. Efficient sulfate production associated with haze formation has been reported in several studies conducted in China (Cheng et al., 2016; Xie et al., 2015; Xue et al., 2016). Ammonium production during haze formation is tied with neutralization of acidic aerosol by ammonia, which was apparently present abundantly in the gas phase. Compared with the EC concentration, light extinction by (NH_4NO_3) increased the most from the highest MLH conditions ($248 M m^{-1}/\mu g m^{-3}$) to the lowest MLH conditions ($1150 M m^{-1}/\mu g m^{-3}$) as shown by Figure 7b. Overall, the rapid growth of nitrate aerosol mass, together with abundant concentration of organic aerosol, were the main cause of the light extinction for dry aerosol under haze formation.



363 The mechanism governing the aerosol-chemistry-radiation-boundary layer feedback for the rapid
364 growth of atmospheric aerosol is illustrated in Fig. 8. As a result of reduction in solar radiation and
365 atmospheric heating, a variety of chemical reactions in the gas phase and on particle surfaces or inside
366 them are enhanced with an increased relative humidity and AWC. Such conditions are unfavorable
367 for the dispersion of pollutants, which further enhances atmospheric stability. The formation of
368 hydrophilic compounds, e.g., nitrate, sulfate and oxygenated organic compounds, result in enhanced
369 water uptake by aerosol particles, which will essentially increase heterogeneous reactions associated
370 with these particles. As a result, the aerosol mass and size increase, light extinction is enhanced, and
371 the development of the mixing layer is depressed. At the same time, aerosol precursors concentrated
372 within a shallower mixing layer lead to enhanced production rate of aerosol components in both gas
373 and aerosol phases, especially nitrate but also other secondary aerosol. The increased concentrations
374 of aerosol will further enhance this positive loop.

375

376 4. Conclusions

377

378 We investigated the synergetic variations of aerosol chemical composition and mixing layer height
379 during the daytime in urban Beijing. Significant dependency of the sharp increase of ammonium
380 nitrate and aerosol water content with the occurrence of the explosive aerosol mass growth events
381 were observed. We showed that these two components drove a positive aerosol-chemistry-radiation-
382 boundary layer feedback loop, which played an important role in the explosive aerosol mass growth
383 events. A plausible explanation is that the increased aerosol water content at low mixing layer heights
384 provides favorable conditions for heterogeneous reactions for nitrate and sulfate production and
385 neutralization by ammonia. The significant formation of secondary aerosol increases the
386 concentration of aerosol particles in the diameter range 300-700 nm, which effectively reduces the
387 solar radiation reaching the surface and further enhances the aerosol-chemistry-radiation-boundary
388 layer feedback loop. Our analysis connects the aerosol light extinction to a reduction in the mixing
389 layer height, which suppresses the volume into which air pollutants are emitted and leads to an



390 explosive aerosol mass growth. Our results indicate that reduction of ammonium and nitrate
391 concentration in aerosol could weaken the aerosol-radiation-chemistry-boundary layer feedback loop,
392 which could thereby reduce heavy haze episodes in Beijing.

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397 6. Competing financial interests

398 The authors declare no competing financial interests.

7. Author contributions

399 YW and MK initiated the study. ZL, YW, FZ, YZ, YG, ZF, CL, YZ, TC, CY, KD, BC, JK, LY, XF,
400 WD, JC and YL conducted the longtime measurements. ZL, YW, LD, RC, SH, PZ, LW, VK, YL and
401 MK interpreted the data. ZL, YW and VK wrote the manuscript.

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Figure caption

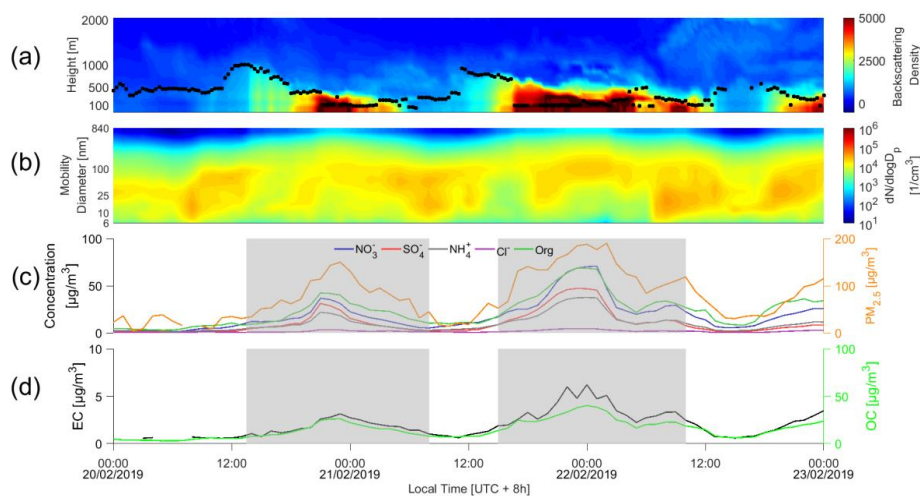


Figure 1. Time series of (a) backscattering density and boundary layer height (b) particle number concentration distribution (PNSD), (c) chemical composition species mass concentration and $PM_{2.5}$ mass concentrations and (d) elemental carbon (EC) and organic carbon (OC). The haze periods are marked by the shaded areas.

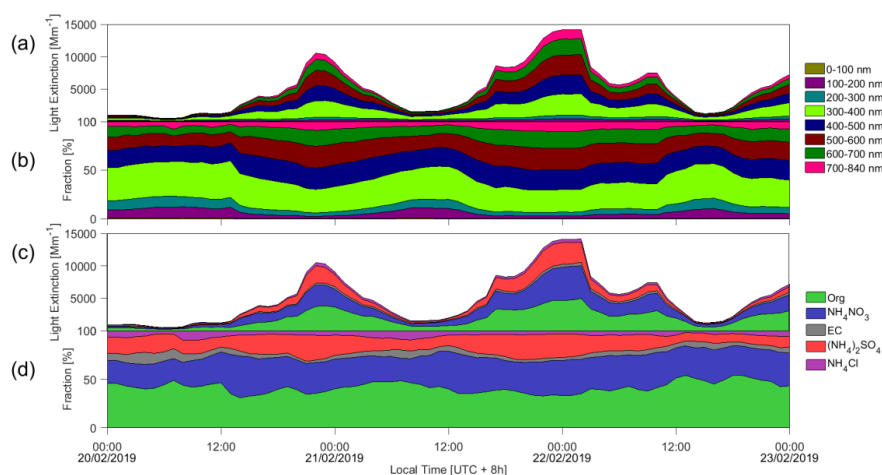
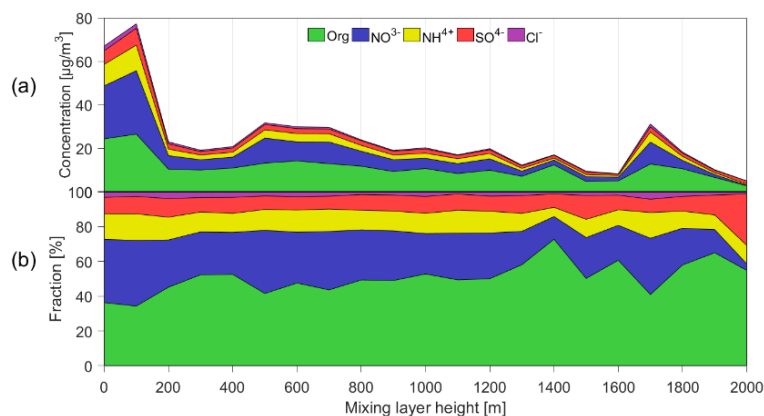


Figure 2. Time series of (a, b) variation of light extinction from different size aerosol and fractions, and (c, d) variation of light extinction from different aerosol species and fractions. The legends in the left side of figures are particle diameter and chemical compositions, respectively.



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 469 Figure 3. Statistical relationship between MLH and concentration (a) and fraction (b) of chemical
 470 composition species. Only daytime conditions determined by ceilometer from non-rainy periods
 471 (RH<95%) are considered.

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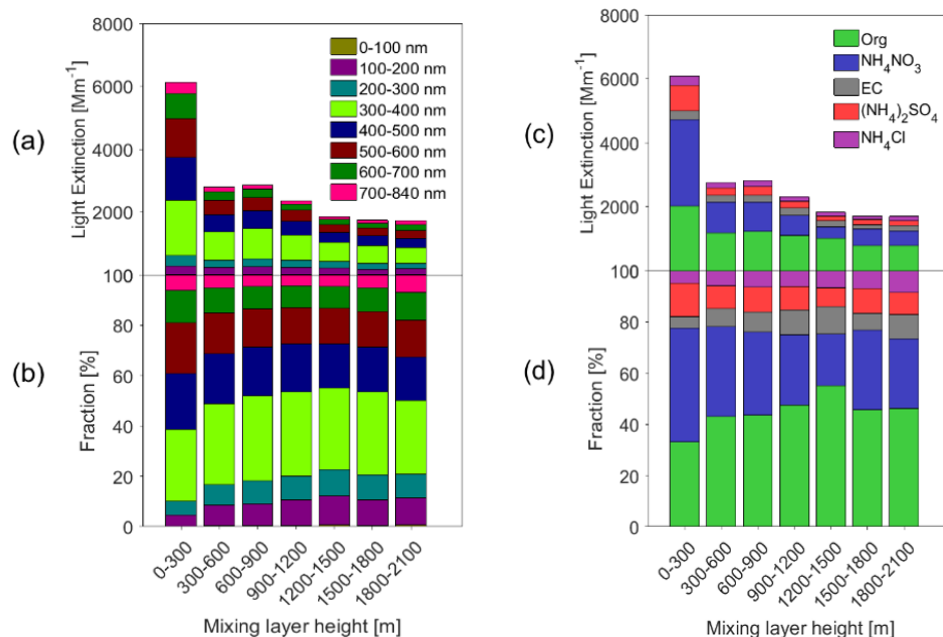
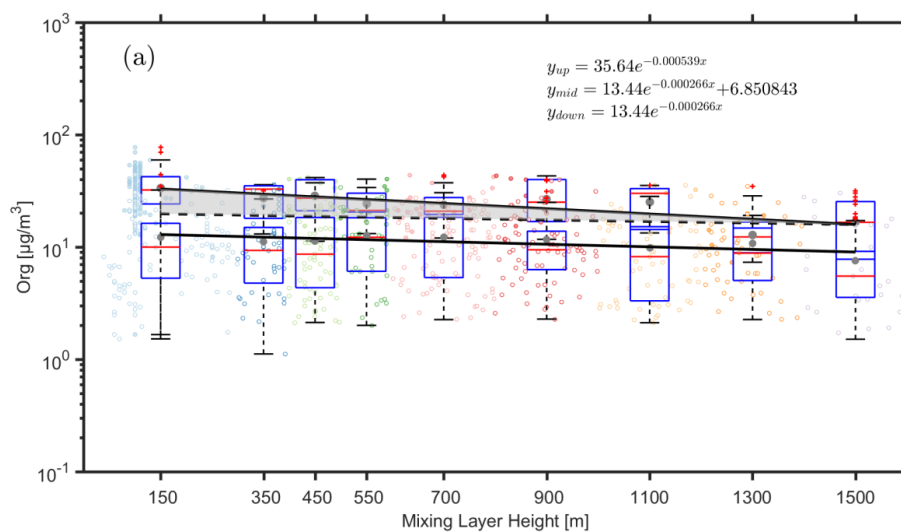
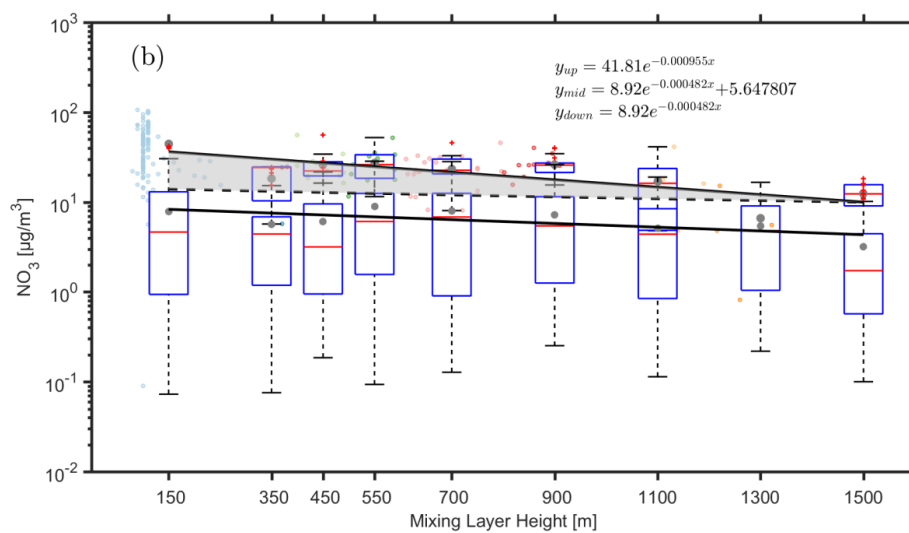


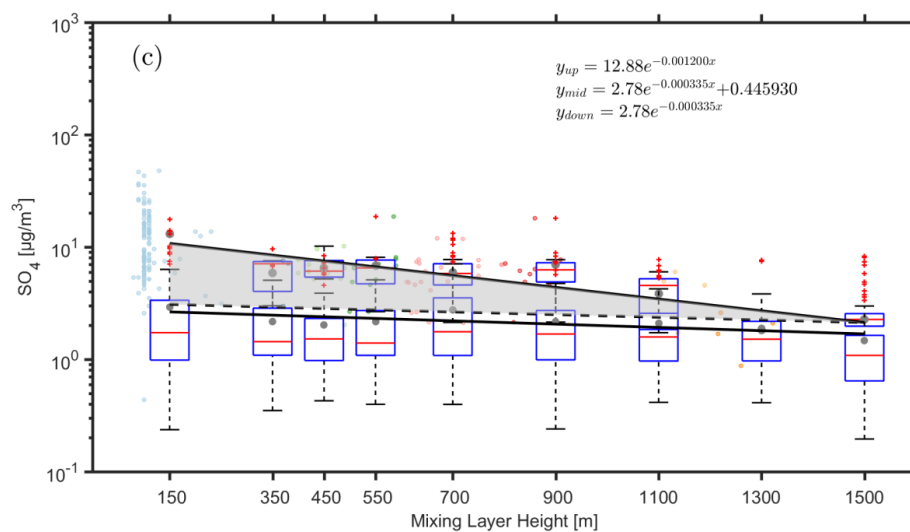
Figure 4. Statistical relationship between MLH and light extinction of different aerosol species. Only daytime conditions determined by the ceilometer from non-rainy periods ($\text{RH} < 95\%$) are considered.



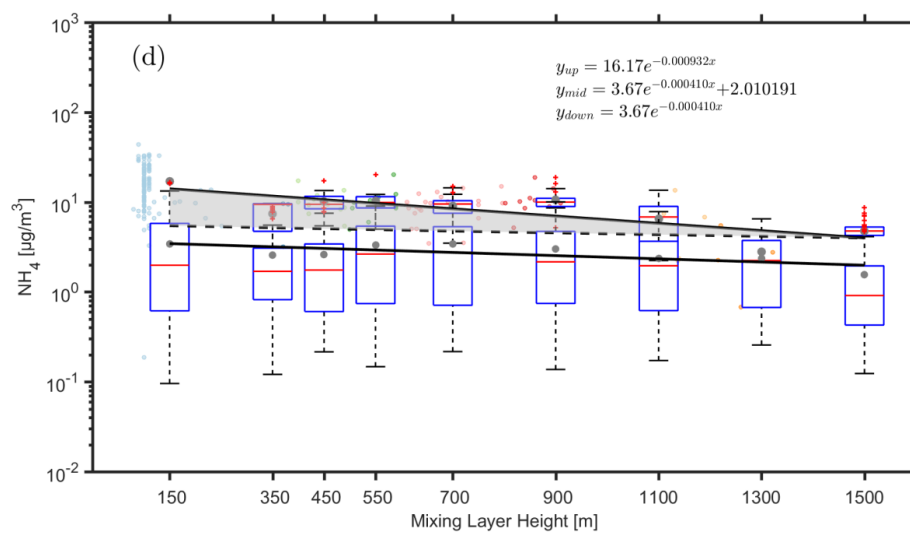
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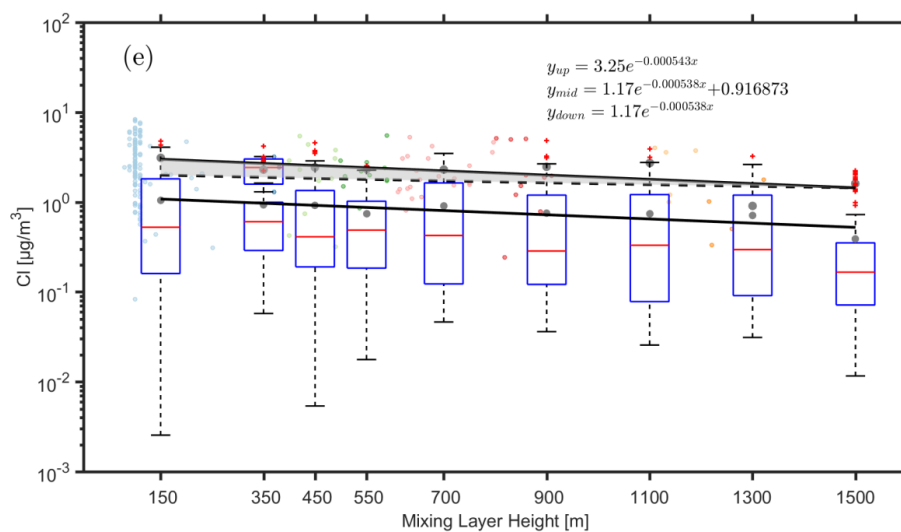
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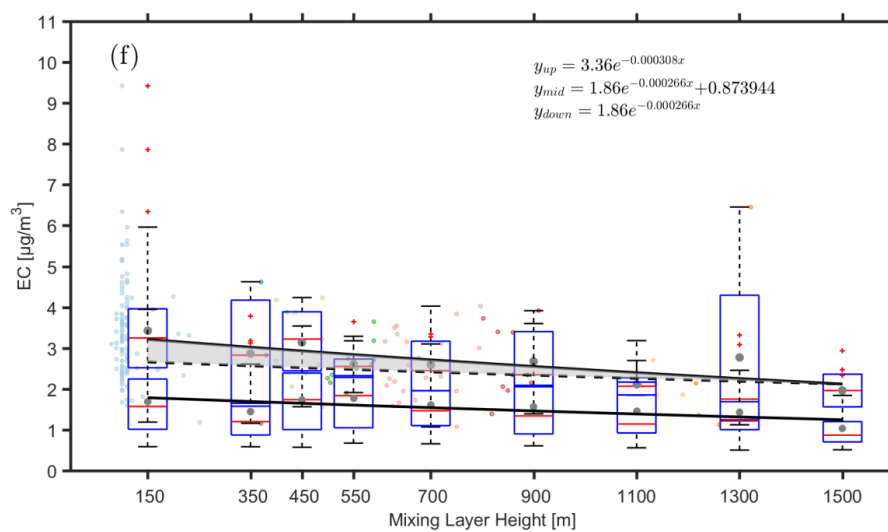
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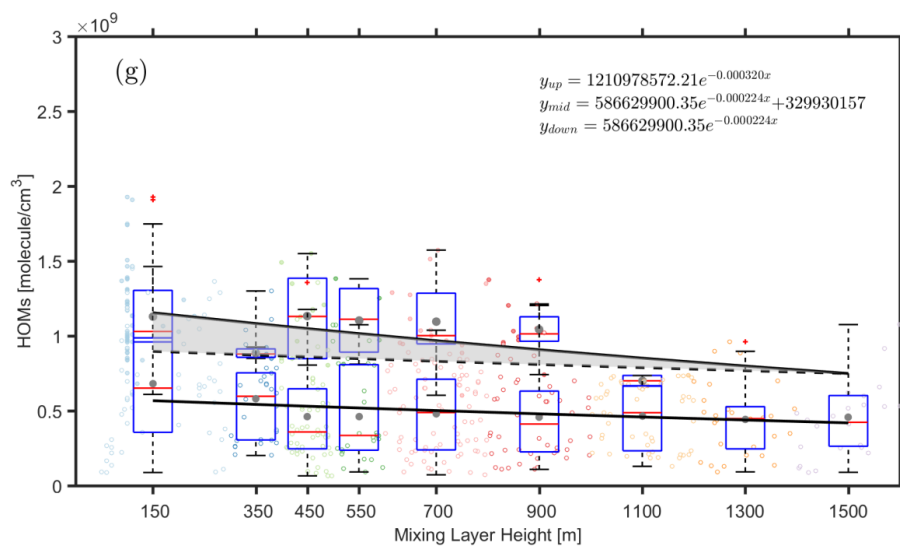
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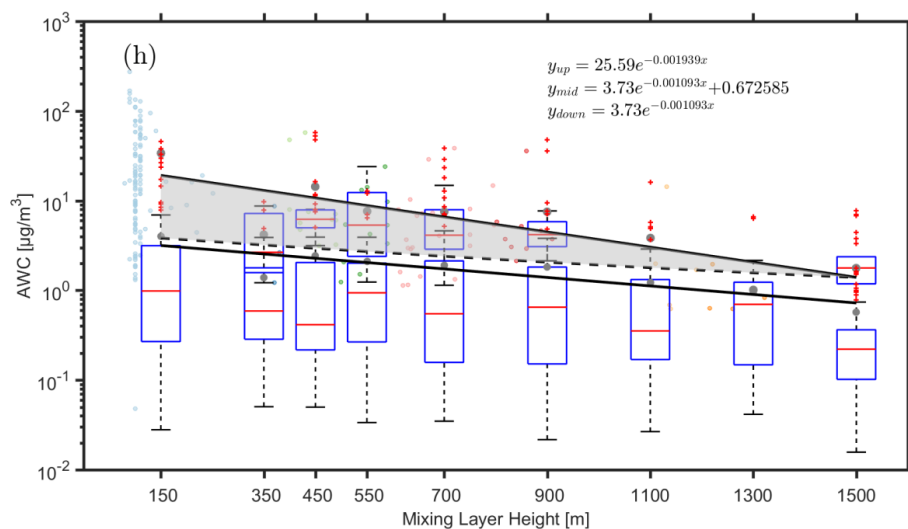


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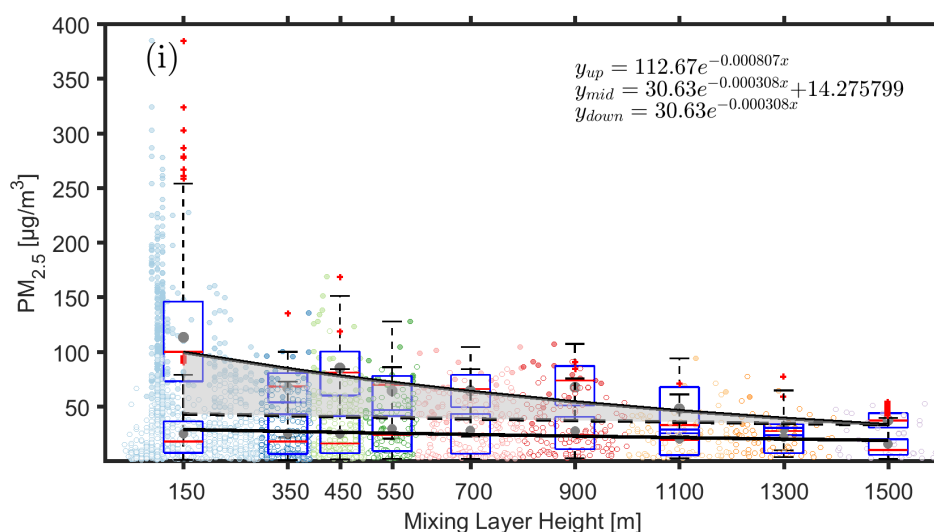
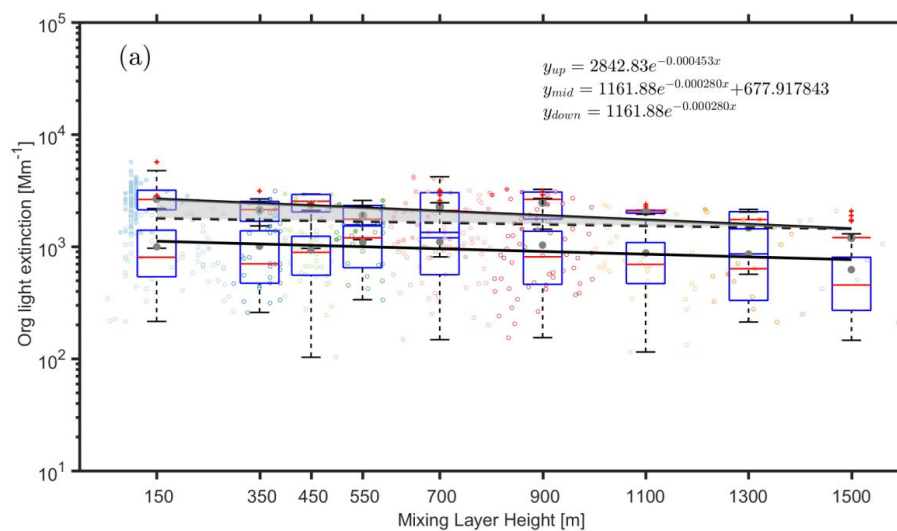


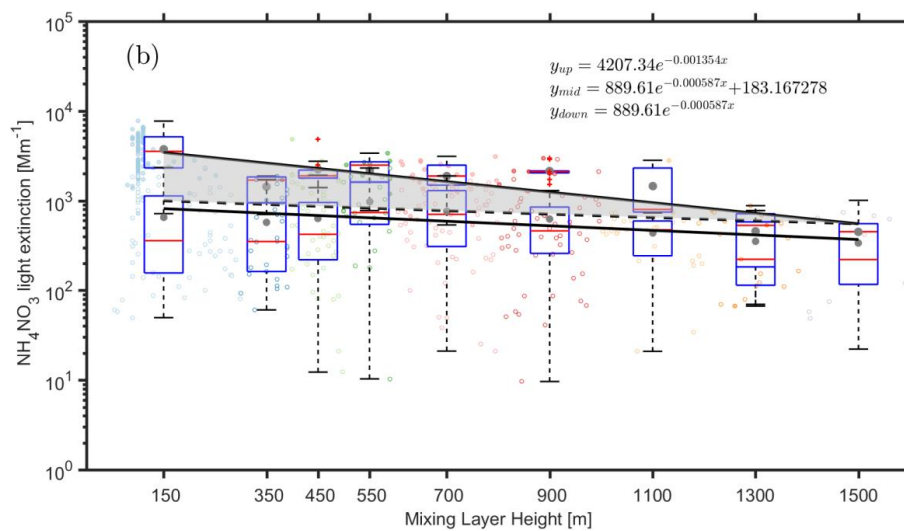
Figure 5. Observed dependency of (organics (a), nitrate (b), ammonium (c), sulfate (d), chlorine (e), element carbon (f), HOMs (g), AWC (h) and PM_{2.5}(i) on the MLH during polluted and less-polluted conditions. The data related to the upper fitting line represents PM_{2.5} concentrations larger than 75 µg m⁻³, while the data related to the lower fitting line represents PM_{2.5} concentrations lower than 75 µg m⁻³. Only daytime conditions determined by the ceilometer from non-rainy periods (RH<95%) were considered. The dark grey points and red lines in the boxes represent mean and median values, respectively. The shaded area between the upper solid and dotted lines corresponds to an increased amount of the specific compounds with decreased MLH, assuming that the compound has the same variation pattern under highly- polluted conditions as in less polluted time.



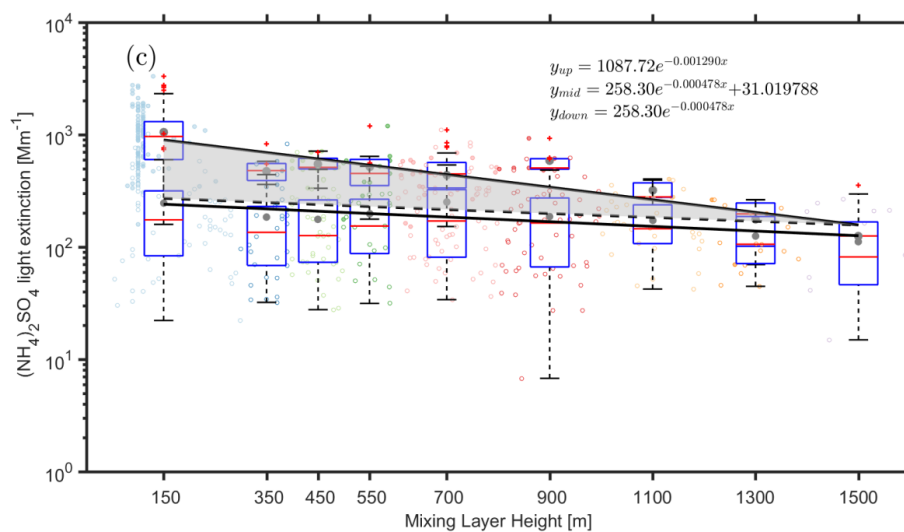
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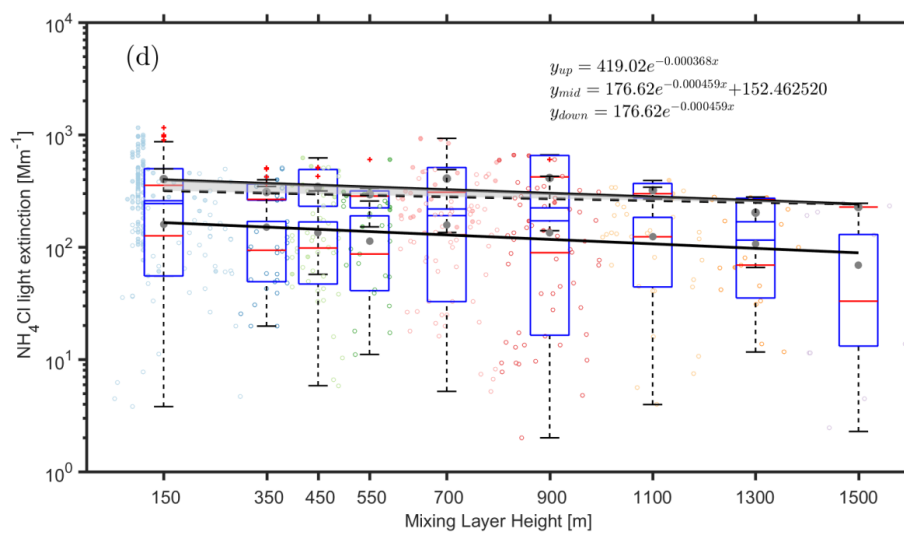
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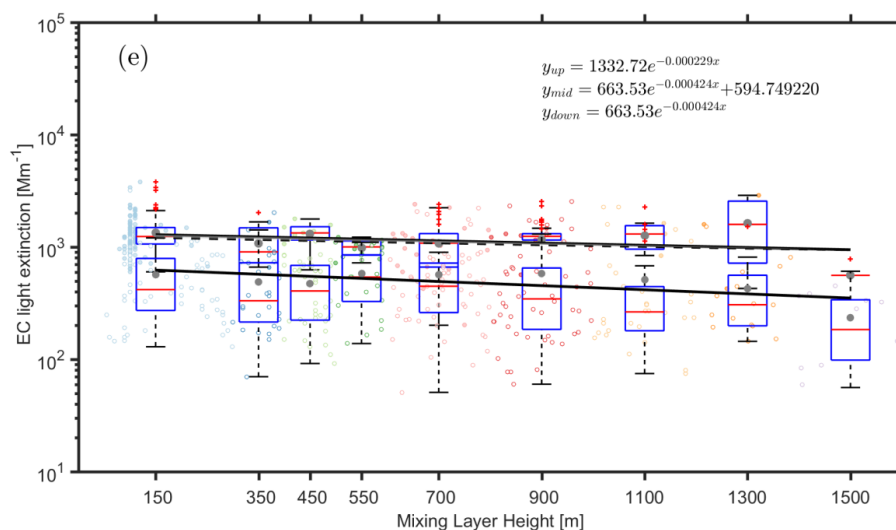


Figure 6. Observed dependency of the aerosol light extinction due to NH_4NO_3 (a) $(\text{NH}_4)_2\text{SO}_4$ (b), NH_4Cl (c) Org (d) and EC (e) on the MLH during polluted and non-polluted conditions. The data related to the upper fitting line represents $\text{PM}_{2.5}$ concentrations larger than $75 \mu\text{g m}^{-3}$, while the data related to the lower fitting line represents $\text{PM}_{2.5}$ concentrations less than $75 \mu\text{g m}^{-3}$. Only daytime conditions determined by ceilometer from non-rainy periods ($\text{RH} < 95\%$) are considered. The dark grey points and red lines in the boxes represent mean and median values, respectively. The shaded area between the upper solid and dashed line corresponds to an increased amount of $\text{PM}_{2.5}$ with a decreased MLH, assuming that $\text{PM}_{2.5}$ has the same variation pattern under highly-polluted conditions as in less polluted time

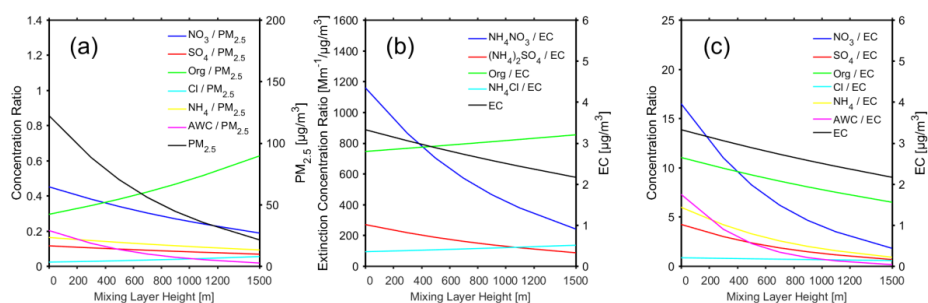


Figure 7. (a) the ratio of the mass concentration of different chemical components (nitrate, sulfate, organics, chlorine, ammonium) and AWC to the mass concentration of $NR_PM_{2.5}$ as a function of MLH. (b) the ratio of dry aerosol light extinction by different chemical components (NH_4NO_3 , $(NH_4)_2SO_4$, Org, NH_4Cl) to the mass concentration EC as a function of MLH (c) the ratio of the mass concentration of different chemical components (nitrate, sulfate, organics, chlorine, ammonium) and AWC to the mass concentration of EC as a function of MLH. All the data corresponds to polluted conditions (fine PM $>75 \mu g m^{-3}$), and only daytime conditions determined by the ceilometer from non-rainy periods ($RH < 95\%$) were considered.

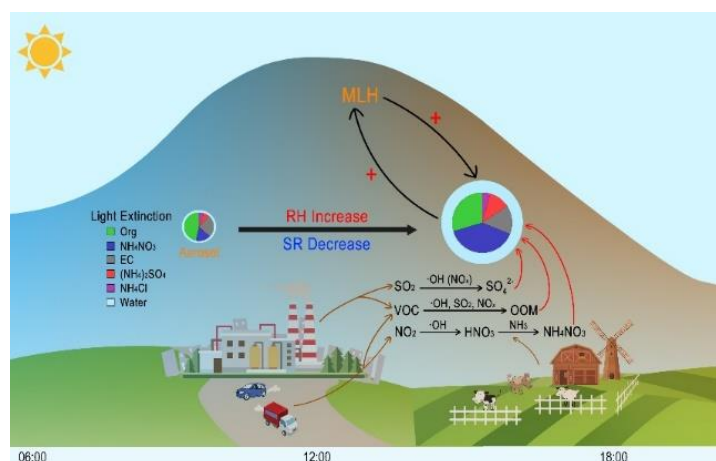


Figure 8. A schematic picture illustrating the process of rapid aerosol mass growth and enhanced light extinction in Beijing. The plus symbols represent the strengthening of a specific process. At the presence of aerosols during afternoon time in Beijing, the intensity of solar radiation reaching the surface will be decreased and relative humidity will be increased. As a result, the development of boundary layer will be suppressed, and the concentrations of aerosol precursors (e.g., SO_2 , NO_2 , VOC) will be increased. In turn, the secondary production of these sulfate, nitrate and oxygenated organic compounds will be enhanced due to increased concentrations and partitioning of these compounds into the aerosol phase. The increased formation of secondary aerosol mass will reduce solar radiation further and the haze formation increased. Noting that during intensive haze periods, nitrate and its contribution to light extinction contribution increased dramatically.