The authors thank the editor and referees to review our manuscript and particularly for the valuable comments and suggestions that are very helpful in improving the manuscript. We provide below point-by-point responses to those comments. We also have made most of the changes suggested by the referees in the revised manuscript.

## Referee #1

In this study, Huang et al. used an aerosol chemical species monitor (ACSM) and an aethalometer to characterize the organic aerosol (OA) in wintertime Beijing. Positive matrix factorization (PMF) was applied to resolve the sources and processes for OA. The effect of RH on the mass concentration, the mass fraction, and the growth rate of various components in PM1 was analyzed. OA was found to dominate the components under both high-RH and low-RH pollution periods. But the change of sulfate and nitrate showed opposite RH-dependence. The results demonstrated the importance of photochemical oxidation and aqueous-phase processes on the formation of secondary aerosol during haze episodes. It could be helpful for understanding the haze formation in wintertime Beijing. Overall, the results are well presented. I recommend the manuscript be considered for publication after following comments being fully addressed.

(1) First of all, the motivation of research should be explained more clearly in the introduction section.

**Response:** We thank the referee's suggestion. In the revised manuscript lines 103-107, we have now added the following description: "Despite the observations of large production of secondary aerosol during haze events, the formation mechanisms are not yet well understood. Specifically, more studies are needed to elucidate the relative importance of photochemical oxidation versus aqueous-phase processes on the formation of secondary aerosol during wintertime haze episodes of different meteorological conditions".

(2) The criterions to distinguish clean day and polluted days, as well as low RH and high RH seem a little bit arbitrary. The selection of the concentration threshold of PM1 for discriminating the clean and pollution period, as well as RH, should be explained.

**Response:** As shown in Figure 1, the clean and pollution episodes occurred alternately during the measurement period, and the PM<sub>1</sub> concentration was usually lower than 20 µg m<sup>-3</sup> during clean episodes and higher than 100 µg m<sup>-3</sup> during polluted episodes. Thus, we divided the measurements into clean period and pollution period using the criteria of 20 µg m<sup>-3</sup> and 100 µg m<sup>-3</sup>, respectively. Furthermore, during the pollution period, RH varied from 15% to 95% with an average value of 46% ( $\approx$  50%) and a median value of 43%, thus we used 50% as the criterion to further divide the pollution period into low-RH pollution days (RH <50%) and high-RH pollution days (RH <50%).

In lines 213-215 in the revised manuscript, we have now added "The clean and pollution episodes occurred alternately during the measurement period, and the  $PM_1$  concentration was usually lower than 20  $\mu$ g m<sup>-3</sup> during clean episodes and higher than 100  $\mu$ g m<sup>-3</sup> during pollution episodes. As such...".

And in lines 220-221, we have added "During the polluted period, RH varied from 15% to 95% with an average value of 46% and a median value of 43%. To investigate...".

(3) It was noted that in Figure 4 where OOA reached a peak value at about 20:00 LT. The authors may give an interpretation about this phenomenon. Also, what is the reason that OOA increased more quickly under low RH condition than under high RH conditions? **Response:** In the section 3.2, we discuss that the diurnal cycle of OOA shows an increase from about 6:00 to 20:00 LT, indicating the contribution from photochemical production. The peak value at about 20:00 LT may be due to the accumulation of OOA formed from photochemical production. Meanwhile, as discussed in section 3.4, photochemical production contributed dominantly to OOA. The O<sub>x</sub> concentration during low-RH pollution days (59.8 ppb) was higher than that during high-RH pollution days (47.8 ppb) and clean days (39.2 ppb). With the higher O<sub>x</sub> concentration (as a surrogate of oxidant level) under low-RH conditions, the daytime formation of OOA was more efficient and the growth rate was higher during those low-RH pollution days than those during high-RH pollution days and clean days.

In line 328 of the revised manuscript, we added "...photochemical production and accumulation of OOA".

(4) Lines 328-332: Why the concentration of Ox (=O3+NO2) was lower than that of NO2? Please check carefully. Meanwhile, the mass fraction of some species (e.g., Cl-) was missing in Fig. 5.

**Response:** Apologies for the typos. We double check our data, and it now reads: " $NO_2$  increased accordingly from 16.7 ppb during clean periods to 42.2 ppb during high-RH pollution periods and to 55.4 ppb during low-RH pollution periods". We have changed the  $NO_2$  values in Table1 accordingly. The missing mass fractions for some species are also added in Figure 5.

(5) Lines 335-337: I don't think there was obvious difference in the mass fraction of nitrate in two types of pollution periods (14 % vs. 15 %). And references need to be cited in Line 337. In addition, this sentence is too hard to read. Rewrite it!

**Response:** We thank the referee for pointing this out. We have made the change and it now reads "We observed similar contributions from nitrate during low-RH pollution periods and high-RH pollution periods, while a much larger contribution from sulfate during high-RH pollution periods than during low-RH pollution periods because of enhanced formation from aqueous-phase processes".

Meanwhile, references were added in line 363, and it now reads "…indicating the importance of residential coal combustion emissions during haze pollution in wintertime Beijing (Elser et al., 2016; Li et al., 2017).

(6) Lines 391-395: Was P3 assigned to high-RH or low-RH pollution during the analyses of sulfate and OA?

Response: P3 was assigned to high-RH pollution episode during the analyses of sulfate

and OOA. We made change in lines 415-416 and it now reads "OOA also showed a much higher contribution to OA during high-RH pollution events (62% for P6 and 50% for P7) than during low-RH pollution events (P1, P2, P4 and P5, 20-31%)".

(7) Line 423. The measurement of SO2 was not described in the Section 2.1 Site description and instrumentation.

**Response:** Thanks for the referee's reminder. In the revised manuscript lines 136-138, we have now added "The gaseous species including  $O_3$ ,  $NO_x$ , and  $SO_2$  were measured by a Thermo Scientific Model 49i ozone analyzer, a Thermo Scientific Model 42i  $NO-NO_2-NO_x$  analyzer, and an Ecotech EC 9850 sulfur dioxide analyzer, respectively".

(8) Conclusion. I think a brief description of atmospheric implications should be included in this section.

**Response:** In the revised manuscript (lines 478-481), we have now added the following discussion "These results provide insights into the relative importance of photochemical oxidation and aqueous-phase processes for secondary aerosol formation during haze pollution, demonstrating the significance of meteorological conditions in the formation of secondary aerosol."

#### Referee #2

This manuscript presents the comparisons of PM1 species and organic aerosol (OA) sources/processes during winter in Beijing for clean and pollution periods, with a particular focus on the effect of relative humidity (RH) on secondary aerosol formation. The comparisons were made through mass concentration, mass fraction, and growth rate. It is found that OA dominated the PM1 mass under both low-RH and high-RH pollution conditions. However, sulfate was found to increase during high-RH pollution periods and nitrate increased during low-RH pollution periods. Oxygenated OA (OOA) showed higher growth rate during low-RH pollution period than during high-RH pollution period. These results provide insights into the relative importance of photochemical oxidation vs. aqueous-phase processes for secondary aerosol formation under different meteorological conditions. It is a useful addition to the literature for understanding the haze formation in Beijing. The manuscript is well written, and results are discussed logically. I recommend publication in ACP after a few minor points are addressed.

(1) In section 2.2.3, organics was not considered in the ALWC calculation using the ISORROPIA-II model. Please provide an explanation.

**Response:** We thank the reviewer for pointing this out. We calculated the ALWC using the ISORROPIA-II model, which simulates the thermodynamic equilibrium of the  $NH_4^+$ - $SO_4^{2-}$ - $NO_3^-$ - $Cl^-$ - $H_2O$  system and does not consider the organics contribution. To evaluate this uncertainty, we further calculate the contribution of organics to ALWC following the

approach in Guo et al (2015) and Cheng et al (2016):

$$W_{org} = \frac{OM}{\rho_{org}} \cdot \rho_{w} \cdot \frac{\kappa_{org}}{(100\%/RH - 1)}$$

where OM is the mass concentration of organics,  $\rho_w$  is the density of water and  $\rho_{org}$  is the density of organics ( $\rho_{org} = 1.4 \times 10^3$  kg m<sup>-3</sup>, Cerully et al., 2014).  $\kappa_{org}$  is the hygroscopicity parameter of organic aerosol composition. We adopted a  $\kappa_{org}$  value of 0.06 based on previous cloud condensation nuclei measurements in Beijing (Gunthe et al., 2011).

The calculated results showed that the average contribution of organics to the total ALWC was 18%, suggesting that inorganic species were the dominant hygroscopic species and organics had a minor contribution to ALWC.

In the revised manuscript, we added the following description in section 2.3.3: "Meanwhile, the contribution of organics to ALWC (ALWC<sub>0</sub>) was also calculated using the following equation (Guo et al., 2015; Cheng et al., 2016):

$$W_{org} = \frac{OM}{\rho_{org}}. \rho_{w}. \frac{\kappa_{org}}{(100\%/RH - 1)}$$

where OM is the mass concentration of organics,  $\rho_w$  is the density of water and  $\rho_{org}$  is the density of organics ( $\rho_{org} = 1.4 \times 10^3$  kg m<sup>-3</sup>, Cerully et al., 2014).  $\kappa_{org}$  is the hygroscopicity parameter of organic aerosol composition. We adopted a  $\kappa_{org}$  value of 0.06 based on previous cloud condensation nuclei measurements in Beijing (Gunthe et al., 2011)." Figure 7 was also updated accordingly.

(2) Page 6, line 203-205, the authors divided the pollution period into low-RH pollution days (RH <50%) and high-RH pollution days (RH >50%). What is the criterion for this definition of low- and high-RH?

**Response:** As discussed in response to referee #1, during the polluted period, RH varied from 15% to 95% with an average value of 46% ( $\approx$  50%) and a median value of 43%. Thus, we used 50% as the criterion to further divide the pollution period into low-RH pollution days (RH <50%) and high-RH pollution days (RH >50%). If 60% is used as a cutting point, the data points (78) in RH >60% are much less than those in RH<60% (282), which may be not proper for statistical comparison.

In the revised manuscript lines 220-221, we have added "During the polluted period, RH varied from 15% to 95% with an average value of 46% and a median value of 43%. To investigated...".

(3) Page 8, line 303, change "OOA is correlated well with nitrate (R2=0.89). and the diurnal cycle : : :: : : " to "OOA is correlated well with nitrate (R2=0.89), and the diurnal cycle : : :: : :".

**Response:** Change made.

(4) Page 9, line 330-334. ": : :: :: We observed a much larger contribution from nitrate

during low-RH pollution periods than during high-RH pollution periods: : :: : : " This is not well supported as both the mass concentration and fraction of nitrate are similar. Please check it carefully.

**Response:** We thank the referee for pointing this out. In the revised manuscript lines 353-358, it now reads "We observed similar contributions from nitrate during low-RH pollution periods and high-RH pollution periods, while a much larger contribution from sulfate during high-RH pollution periods than during low-RH pollution periods because of enhanced formation from aqueous-phase processes".

(5) In Table 1 and throughout the manuscript, the authors should pay attention to the significant digits which denotes precision of measurements.

**Response:** Thanks. In the revised manuscript, we have now unified the significant digits throughout the manuscript and Table 1.

#### Referee #3

This paper focuses on discussing the secondary aerosol formation processes under different RH conditions mainly. It is valuable to understand the aerosol chemistry in Beijing, but I have a few serious concerns on current results and interpretation. See below:

(1) One concern is that the dataset used is relatively old (five years ago). It is therefore not very up-to-date to reflect the real processes in current atmosphere given the concentrations, compositions of PM1 as well as the precursors might have changed greatly in Beijing. The authors have to comment more on the implications of findings here. **Response:** We agree with the referee that the concentrations, compositions of  $PM_1$  as well as the precursors might have some changes in recent years. Our measurements were conducted after the implementation of legislative 'Air Pollution Prevention and Control Action Plan' in 2013. Therefore, our results show some similar variations on the  $PM_1$ composition when compared to more recent studies, such as the increase of nitrate and decrease of sulfate. Meanwhile, our study focused on the formation mechanisms of secondary aerosol during different meteorological conditions in haze pollution, especially the significant effects of RH, which is still not well elucidated. The results demonstrated the importance of photochemical oxidation and aqueous-phase processes on the formation of secondary aerosol during haze episodes. It could be helpful for understanding the haze formation in wintertime Beijing. Therefore, our study still provides valuable information to the scientific community to improve our understanding of fine PM pollution.

In the revised manuscript, we have now added the following discussion in conclusion: "These results provide insights into the relative importance of photochemical oxidation and aqueous-phase processes for secondary aerosol formation during haze pollution, demonstrating the significance of meteorological conditions in determining the formation of secondary aerosol". (2) More details regarding the PMF analyses should be provided. It is not clear why 5-factor solution is optimal, considering that you use ACSM data which had very limited chemical resolution to identify tracer ions. And you used ME-2 technique, which on one hand is better to extract the real factors presenting in your data, but on the other hand, you may presume and artificially identify a factor that might not be real. Justification of the PMF results is essential and why and how the initial profiles of different factors were used are not clear. I feel that current information provided here is not enough.

Response: Thanks for the referee's suggestion. We first performed free PMF (unconstrained PMF) runs to identify the main factors. The number of factors was determined by examining three to seven factors, and the factors were identified based on their mass spectral profiles and correlation with external tracers. The five-factor solution (i.e., HOA, COA, CCOA, BBOA, and OOA) was selected because it best explains our data. For the solutions with less than 5 factors, HOA appeared to be mixed with COA while CCOA mixed with BBOA (Fig. S1). However, when the number of factors was increased to 6, the 00A factor split into two 00A factors of similar time series (Fig. S2), suggesting that further separation of the factors does not improve the interpretation of the data. After determining the number of factors (five) using un-constrained PMF runs, we then performed ME-2 runs, which explore the variabilities of certain (mainly primary) factors (i.e., HOA, COA, and BBOA) by constraining their profiles with *a* value approach. In a word, the un-constraint PMF helps making sure that the five factors were real (without "forced deconvolution" as in constrained ME-2), while the constrained ME-2 helps refining the contributions from those five factors. Please see Fig. S1-3 and the corresponding discussion in Section 3.2 for further details of the PMF/ME-2.

(3) Calculation of ALWC by using ISORROPIA-II model had uncertainties as it only consider water uptake by inorganic species but organics is dominant your PM1, please comment on this, and discuss the influences on your results.

**Response:** We thank the reviewer for pointing this out. We calculated the ALWC using the ISORROPIA-II model, which simulates the thermodynamic equilibrium of the  $NH_4^+$ – $SO_4^2-NO_3^-$ – $Cl^-$ – $H_2O$  system and does not consider the organics contribution. To evaluate this uncertainty, we further calculate the contribution of organics to ALWC following the approach in Guo et al (2015) and Cheng et al (2016):

$$W_{org} = \frac{OM}{\rho_{org}} \cdot \rho_{w} \cdot \frac{\kappa_{org}}{(100\%/RH - 1)}$$

where OM is the mass concentration of organics,  $\rho_w$  is the density of water and  $\rho_{org}$  is the density of organics ( $\rho_{org} = 1.4 \times 10^3$  kg m<sup>-3</sup>, Cerully et al., 2014).  $\kappa_{org}$  is the hygroscopicity parameter of organic aerosol composition. We adopted a  $\kappa_{org}$  value of 0.06 based on previous cloud condensation nuclei measurements in Beijing (Gunthe et al., 2011).

The calculated results showed the average contribution of organics to the total ALWC was 18%, suggesting that inorganic species were the dominant hygroscopic species and organics had minor contribution to ALWC.

In the revised manuscript, we have added following description in section 2.3.3: "Meanwhile, the contribution of organics to ALWC(ALWC<sub>0</sub>) was also calculated using the following equation (Guo et al., 2015; Cheng et al., 2016):

$$W_{\text{org}} = \frac{OM}{\rho_{\text{org}}} \cdot \rho_{\text{w}} \cdot \frac{\kappa_{\text{org}}}{(100\%/\text{RH} - 1)}$$

where OM is the mass concentration of organics,  $\rho_w$  is the density of water and  $\rho_{org}$  is the density of organics ( $\rho_{org} = 1.4 \times 10^3$  kg m<sup>-3</sup>, Cerully et al., 2014).  $\kappa_{org}$  is the hygroscopicity parameter of organic aerosol compositions. We adopted a  $\kappa_{org}$  value of 0.06 based on previous cloud condensation nuclei measurements in Beijing (Gunthe et al., 2011)." Also Figure 7 was updated accordingly.

(4) A very serious concern is the large uncertainty of your interpretation. In order to make strong argument regarding the different chemical processes under different RH conditions. You have to eliminate the influences of other meteorological conditions (PBL, wind directions, speeds, and different air masses) on the concentrations, compositions and growth rates you investigated here. Otherwise, you cannot claim that the observed changes were solely due to chemistry. This reviewer see very little discussions regarding this point, and this make the results highly untrustworthy. For example, the calculation of growth rates, such rates is largely not due to chemistry but likely PBL variations, etc. In understand that the authors argue that during pollution period, there was low wind and mainly south/southeasterly wind mainly; this is too general and does not help resolve what I mention here.

**Response:** We agree with the referee that factors, such as PBL height and winds, might complicate the interpretation of chemical processes under different RH conditions. For example, the growth rates of different species might be affected by those factors. Here in this study, however, we did not intend to emphasize on the absolute values of growth rates, which make less sense given the factors mentioned above. Instead, we mostly tried to compare the growth rates of different species (e.g., sulfate and nitrate in lines 243-244) under the same periods, which should share the effects by those factors. When we did compare the growth rates of the same species under different time periods (low-RH and high-RH pollution conditions), as also pointed out by the reviewer, we assumed that the influences from those factors were more or less similar for these two types of pollution conditions. To provide a note of caution, the caveat of this kind of comparison is now added in lines 250-255 in the revised manuscript: "Note that the comparison of growth rates was done under the assumption that chemical processes were the main reason for mass growth, which might not be the case if other factors such as planetary boundary layer height variations dominate. Yet comparison of growth rates of different species in the same time period would not be affected by these factors because those species should share the same effects".

(5) L199-L201: Table 1 does not provide wind directions as you said. **Response:** Apologies for the typo. Table 1 has been changed to Fig. 1.

(6) Indeed, similar discussion had been published in a few references cited here, and it seems to be a bit superficial here, especially section 3.4. The authors need to add more discussions, and point out clearly what are the unique and novel findings here from other studies.

**Response:** We agree with the referee that some similar discussion had been published in a few references cited here. However, the causes of fine PM pollution in urban Beijing are still not fully understood, most likely due to the campaign-to-campaign difference in meteorological conditions, emissions, and atmospheric processes. In addition, previous studies usually focused on sources variations or only SIA or SOA formations (Sun et al., 2015; Hu et al., 2016; Hu et al., 2017; Xu et al., 2017). In our study, we focused on the RH effects during haze pollution. The effect of RH on the mass concentration, the mass fraction, and the growth rate of various components in PM<sub>1</sub> was all analyzed. OA was found to dominate the components under both high-RH and low-RH pollution periods. But the change of sulfate and nitrate showed opposite RH-dependence. The results demonstrated the importance of photochemical oxidation and aqueous-phase processes on the formation of secondary aerosol during haze episodes. It could be helpful for understanding the haze formation in wintertime Beijing. Therefore, our study still provides valuable information to the scientific community to improve our understanding of fine PM pollution.

In the revised manuscript, we have now added the following discussion in section 3.4 in lines 436-438: "Meanwhile, the high ratio between  $NO_{3}$  and  $SO_{4}^{2}$  suggest the nitrate production is more efficient than that of sulfate during low-RH pollution period", and in lines 447-452: "Meanwhile, the  $O_x$  concentration during low-RH pollution days (59.8 ppb) was higher than that during high-RH pollution days (47.8 ppb) and clean days (39.2 ppb). With the higher  $O_x$  concentration (as a surrogate of oxidant level) under low-RH conditions, the daytime formation of OOA was more efficient and the growth rate was higher during those low-RH pollution days than those during high-RH pollution days and clean days".

And added the following discussion in the conclusion in lines 478-481: "These results provide insights into the relative importance of photochemical oxidation and aqueousphase processes for secondary aerosol formation during haze pollution, demonstrating the significance of meteorological conditions in determining the formation of secondary aerosol".

# (7) Why you chose 50% RH as a cutting point for low- and high-RH conditions? How about 60%, and how does this choice possibly influence your findings?

**Response:** During the pollution period, RH varied from 15% to 95% with an average value of 46% and a median value of 43%. Thus, we used 50% as the criterion to further divide the pollution period into low-RH pollution days (RH <50%) and high-RH pollution days (RH >50%). If 60% is used as a cutting point, the data points (78) in RH >60% are much less than those in RH<60% (282), which may be not proper for statistical comparison. Meanwhile, this choice has minor influence on our findings. For example, when using 60% RH as the cutting point, Figure R1 (left) and Figure R2 (right) are very similar to those

two figures (Fig. 6 and Fig. 7) in the manuscript. The better correlation between SOR and ALWC in high RH condition and the conclusion of "Aqueous-phase production of  $SO_4^{2-}$  become important during high RH periods" are still valid. Also, as shown in Figure R1 (left) below, when using 60% RH, some data in 50%~60% shows similar trend with that of RH>60%, further confirming that 50% RH is a better choice as the cutting point for low-and high-RH pollution conditions.



Fig. R1. The relationship between  $SO_{4^{2-}}$  and  $NO_{3^{-}}$  during low-RH (RH <60%) and high-RH (RH >60%) pollution episodes (left).

Fig. R2. The relationship between the sulfate oxidation ratio (SOR =  $[SO_4^{2-}]/([SO_4^{2-}] + [SO_2])$  and ALWC at high RH pollution condition (RH >60%) (right).

In the revised manuscript lines 220-221, we added "During the polluted period, RH varied from 15% to 95% with an average value of 46% and a median value of 43%. To investigate...".

## Reference:

- Cerully, K. M., Bougiatioti, A., Hite Jr., J. R., Guo, H., Xu, L., Ng, N. L., Weber, R., and Nenes, A.: On the link between hygroscopicity, volatility, and oxidation state of ambient and water-soluble aerosols in the southeastern United States, Atmos. Chem. Phys., 15, 8679–8694, https://doi.org/10.5194/acp-15-8679-2015, 2015.
- Cheng, Y. F., Zheng, G. J., Wei, C., Mu, Q., Zheng, B., Wang, Z. B., Gao, M., Zhang, Q., He, K. B., Carmichael, G., Pöschl, U., and Su, H.: Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events in China, Sci. Adv., 2, e1601530, https://doi.org/10.1126/sciadv.1601530, 2016.
- Elser, M., Huang, R. J., Wolf, R., Slowik, J. G., Wang, Q., Canonaco, F., Li, G., Bozzetti, C., Daellenbach, K. R., Huang, Y., Zhang, R., Li, Z., Cao, J., Baltensperger, U., El-Haddad, I., and Prévôt, A. S. H.: New insights into PM2.5 chemical composition and sources in two major cities in China during extreme haze events using aerosol mass spectrometry, Atmos. Chem. Phys., 16, 3207–3225, https://doi.org/10.5194/acp-16-3207-2016, 2016.
- Gunthe, S. S., Rose, D., Su, H., Garland, R. M., Achtert, P., Nowak, A., Wiedensohler, A., Kuwata, M., Takegawa, N., Kondo, Y., Hu, M., Shao, M., Zhu, T., Andreae, M. O., and Poschl, U.:Cloud condensation nuclei (CCN) from fresh and aged air pollution in the megacity region of

Beijing, Atmos. Chem. Phys., 11, 11023–11039, 2011.

- Guo, H., Xu, L., Bougiatioti, A., Cerully, K.M., Capps, S.L., Hite, J.R., Jr, Carlton, A.G., Lee, S.H., Bergin, M.H., Ng, N.L.: Fine-particle water and pH in the southeastern United States, Atmos. Chem. Phys., 15, 5211-5228, 2015.
- Hu, W., Hu, M., Hu, W., Jimenez, J. L., Yuan, B., Chen, W., Wang, M., Wu, Y., Chen, C., Wang, Z., Peng, J., Zeng, L., and Shao, M.: Chemical composition, sources, and aging process of submicron aerosols in Beijing: Contrast between summer and winter, J. Geophys. Res. Atmos., 121(4), 1955–1977, https://doi.org/10.1002/2015JD024020, 2016.
- Hu, W., Hu, M., Hu, W.-W., Zheng, J., Chen, C., Wu, Y., and Guo, S.: Seasonal variations in high time-resolved chemical compositions, sources, and evolution of atmospheric submicron aerosols in the megacity Beijing, Atmos. Chem. Phys., 17, 9979–10000, https://doi.org/10.5194/acp-17-9979-2017, 2017.
- Li, H., Zhang, Q., Zhang, Q., Chen, C., Wang, L., Wei, Z., Zhou, S., Parworth, C., Zheng, B., Canonaco, F., Prévôt, A. S. H., Chen, P., Zhang, H., Wallington, T. J., and He, K.: Wintertime aerosol chemistry and haze evolution in an extremely polluted city of the North China Plain: significant contribution from coal and biomass combustion, Atmos. Chem. Phys., 17, 4751–4768, https://doi.org/10.5194/acp-17-4751-2017, 2017.
- Sun, Y. L., Wang, Z. F., Du, W., Zhang, Q., Wang, Q. Q., Fu, P. Q., Pan, X., Li, J., Jayne, J., and Worsnop, D. R.: Long-term real-time measurements of aerosol particle composition in Beijing, China: seasonal variations, meteorological effects, and source analysis, Atmos. Chem. Phys., 15, 10149–10165, https://doi.org/10.5194/acp-15-10149-2015, 2015.
- Xu, W. Q., Han, T. T., Du, W., Wang, Q. Q., Chen, C., Zhao, J., Zhang, Y. J., Li, J., Fu, P. Q., Wang, Z. F., Worsnop, D. R., and Sun, Y. L.: Effects of Aqueous-Phase and Photochemical Processing on Secondary Organic Aerosol Formation and Evolution in Beijing, China, Environ. Sci. Technol., 51(2), 762–770, https://doi.org/10.1021/acs.est.6b04498, 2017.

# 1 Contrasting sources and processes of particulate species in haze days with low and

- 2 high relative humidity in wintertime Beijing
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- 28

# 29 Abstract

Although there are many studies of particulate matter (PM) pollution in Beijing, the 30 31 sources and processes of secondary PM species during haze periods remain unclear. 32 Limited studies have investigated the PM formation in highly-polluted environments under low and high relative humidity (RH) conditions. Herein, we present a systematic 33 comparison of species in submicron particles ( $PM_1$ ) in wintertime Beijing (29 December 34 35 2014 to 28 February 2015) for clean periods and pollution periods under low and high RH conditions. PM<sub>1</sub> species were measured with an aerosol chemical species monitor 36 (ACSM) and an aethalometer. Sources and processes for organic aerosol (OA) were 37 38 resolved by positive matrix factorization (PMF) with multilinear engine 2 (ME-2). The 39 comparisons for clean, low-RH pollution, and high-RH pollution periods are made from 40 three different aspects, namely (a) mass concentration, (b) mass fraction, and (c) growth rate in diurnal profiles. OA is the dominant component of  $PM_1$  with an average mass 41

concentration of 56.7 µg m<sup>-3</sup> (46%) during high-RH pollution and 67.7 µg m<sup>-3</sup> (54%) 42 43 during low-RH pollution periods. Sulfate had higher concentration and mass fraction during high-RH pollution periods, while nitrate had higher concentration and mass 44 fraction during low-RH pollution periods. The diurnal variations of nitrate and 45 oxygenated organic aerosol (OOA) showed a daytime increase of their concentrations 46 47 during all three types of periods. Nitrate had similar growth rates during low-RH (0.40 µg m<sup>-3</sup> h<sup>-1</sup>) and high-RH (0.55  $\mu$ g m<sup>-3</sup> h<sup>-1</sup>) pollution periods. OOA had a higher growth rate 48 during low-RH pollution periods (1.0 µg m<sup>-3</sup> h<sup>-1</sup>) than during high-RH pollution periods 49 50  $(0.40 \ \mu g \ m^{-3} \ h^{-1})$ . In contrast, sulfate had a decreasing trend during low-RH pollution 51 periods, while it increased significantly with a growth rate of 0.81  $\mu$ g m<sup>-3</sup> h<sup>-1</sup> during high-52 RH pollution periods. These distinctions in mass concentrations, mass fractions, and 53 daytime growth rates may be explained by the difference in the formation processes, 54 affected by meteorological conditions. In particular, photochemical oxidation and aqueous-phase processes may both produce sulfate and nitrate. The relative importance 55 of the two pathways, however, differs under different meteorological conditions. 56 57 Additional OOA formation under high-RH (>70%) conditions suggests aqueous-related 58 formation pathways. This study provides a general picture of the haze formation in Beijing 59 under different meteorological conditions.

## 60 **1 Introduction**

61 Air pollution is a serious environmental problem in China, particularly in the North China Plain (NCP) in winter, affecting air quality and human health. Beijing is one of the most 62 polluted megacities in the NCP, with an annual mean concentration of PM<sub>2.5</sub> being 86 and 63 64  $51 \,\mu g \, m^{-3}$  in 2014 and 2018, respectively (http://sthjj.beijing.gov.cn/), which significantly 65 exceeded the Chinese National Ambient Air Quality Standard (annual average of 35 µg m-<sup>3</sup>). Fine PM pollution in polluted urban environments is complex and is typically 66 67 associated with enhanced primary emissions from multiple sources, strong secondary 68 aerosol formation, and stagnant weather conditions (Sun et al., 2011; 2013; 2016; Huang 69 et al., 2014; Hu et al., 2016; An et al., 2019). Regional transport of air pollutants from 70 urbanized and industrialized areas has an important contribution to fine PM pollution in the NCP region. For example, severe fine PM pollution in Beijing during winter often 71 72 happened when prevailing air masses were from the south (Sun et al., 2016).

73 Organic aerosol (OA) is the major constituent of fine PM and is much less understood compared to inorganic aerosol in terms of their chemical nature and sources (Hallquist et 74 75 al., 2009; Shrivastava et al., 2017). OA is composed of a wide variety of organic species 76 from different sources, and its emission sources and atmospheric processes are not well 77 understood so far, especially in those regions with high fine PM pollution. OA is either 78 directly emitted to the atmosphere (primary organic aerosol, POA) or formed in the 79 atmosphere (secondary organic aerosol, SOA). Therefore, it is essential to identify and 80 quantify the major emission sources and understand the formation processes of OA.

81 The Aerodyne aerosol chemical speciation monitor (ACSM) with quadrupole (Q) or time-

of-flight (TOF) mass analyzer is capable of real-time determination of non-refractory 82 components in submicron aerosol (NR-PM<sub>1</sub>), overcoming the limitation of filter 83 84 measurements such as limited time resolution or measurement artifacts (Ng et al., 2011a; 85 Froehlich et al., 2013). ACSM has been widely used for fine PM studies in many sites in China including Beijing, Nanjing, Shijiazhuang, and Baoji (Sun et al., 2014; Wang et al., 86 2017; Zhang et al., 2017; Huang et al., 2019). By applying positive matrix factorization 87 (PMF, Paatero et al., 1993) or multilinear engine (ME-2, Canonaco et al., 2013) solver to 88 the ACSM data, main OA sources can be identified. Those sources include hydrocarbon-89 90 like OA (HOA), biomass burning OA (BBOA), cooking OA (COA), coal combustion OA (CCOA) 91 and oxygenated OA (OOA). OOA can further be resolved into semi-volatile OOA (SV-OOA) 92 and low-volatility OOA (LV-OOA) by volatility, or more-oxidized OOA (MO-OOA) and less-93 oxidized OOA (LO-OOA) by oxidation state. MO-OOA and LO-OOA together were found to 94 contribute 61% of OA in Beijing during summer in 2011 (Sun et al., 2012), while POA was 95 found to be more important during winter of the same year (Sun et al., 2013). However, many recent studies show large contributions of SOA in wintertime Beijing (Huang et al., 96 97 2014; Hu et al., 2016; Xu et al., 2018) and CCOA is often found to be a large fraction of POA during wintertime pollution days in Beijing (Sun et al., 2016b; Wang et al., 2015; Elser et 98 99 al., 2016). The discrepancies in SOA contribution in different measurement periods reflect the difference in atmospheric and meteorological conditions, e.g., atmospheric oxidative 100 capacity and relative humidity (RH) (Sun et al., 2013; Xu et al., 2017; Wu et al., 2018; Song 101 et al., 2019). 102 103 Despite the observations of large production of secondary aerosol during haze events, the 104 formation mechanisms are not yet well understood. Specifically, more studies are needed 105 to elucidate the relative importance of photochemical oxidation versus aqueous-phase 106 processes on the formation of secondary aerosol during wintertime haze episodes of 107 different meteorological conditions. In this study, we present measurement results at an urban site in Beijing during the winter of 2014-2015. The chemical nature of NR-PM<sub>1</sub>, 108

sources of OA, formation processing of secondary aerosol in different episodes, andparticularly the effects of RH on secondary aerosol formation are discussed.

# 111 2 Methods

# 112 **2.1 Site description and instrumentation**

The online measurements were conducted on the rooftop of a building (about 20 m above the ground level) at the campus of the National Centre for Nanoscience and Technology (40.00° N, 116.38° E) from 29 December 2014 to 28 February 2015. The observation site is between the 4<sup>th</sup> and 5<sup>th</sup> ring roads in the northwest of Beijing and is surrounded by a residential area.

A Q-ACSM was deployed for the mass concentration measurements of NR-PM<sub>1</sub> species, and the detailed operation principles can be found in Ng et al. (2011a). Briefly, ambient air was pumped through a 3/8 in stainless steel tube at a flow rate of 3 L min<sup>-1</sup>, of which

121 85 mL min<sup>-1</sup> was sampled into the Q-ACSM. In order to remove coarse particles, an URG

cyclone (URG-2000-30ED, size cut-off 2.5 µm) was installed in front of the inlet. Because 122 particle bounce can affect collection efficiency (CE), to reduce this uncertainty and to dry 123 the particles, a Nafion dryer (MD-110-48S; Perma Pure, Inc., Lakewood, NJ, USA) was 124 125 installed after the URG cyclone. An aerodynamic lens was used to focus the submicron particles into a narrow beam, the particles beam then impinged on a heated tungsten 126 surface (about 600 °C) to evaporate, impacted by 70-eV electron to ionize, and then 127 detected by a quadrupole mass spectrometer. During this study, the scan rate of Q-ACSM 128 was at 200 ms amu<sup>-1</sup> from m/z 10 to 150 and the time resolution was 30 min. To determine 129 130 the response factor (RF), a differential mobility analyzer (DMA, TSI model 3080) and a 131 condensation particle counter (CPC, TSI model 3772) were used to select and count the 132 monodisperse 350-nm ammonium nitrate (NH<sub>4</sub>NO<sub>3</sub>) particles, respectively. The mass of NH<sub>4</sub>NO<sub>3</sub> particles was calculated with known particle size and number concentrations. 133 134 This calculated mass concentration was compared to the RF of the Q-ACSM, resulting in the ionization efficiency (IE) value (Ng et al., 2011a). 135

The gaseous species including  $O_3$  and  $_1 NO_{x_1}$  and  $SO_2$  were measured by a Thermo Scientific 136 137 Model 49i ozone analyzer-and, a Thermo Scientific Model 42i NO-NO<sub>2</sub>-NO<sub>x</sub> analyzer, and 138 an Ecotech EC 9850 sulfur dioxide analyzer, respectively. The NH<sub>3</sub> concentrations were measured by an NH<sub>3</sub> analyzer (Picarro G2103). The concentrations of black carbon (BC) 139 was determined by an aethalometer (Model AE-33, Magee Scientific) with a time 140 141 resolution of 1 min. In brief, light attenuation at seven different wavelengths was recorded 142 for particle-laden filter spots, and BC concentration was retrieved based on the light 143 attenuation at 880 nm. An automatic weather station (MAWS201, Vaisala, Vantaa, Finland) was used to measure the meteorological parameters including temperature, pressure, 144 145 relatively humidity and visibility, and a wind sensor (Vaisala Model QMW101-M2) was 146 used to measure the wind speed and wind direction.

# 147 2.2 Data analysis

# 148 2.2.1 ACSM data analysis

149 The standard Q-ACSM data analysis software (v.1.5.3.5) written in Igor Pro (WaveMetrics, 150 Inc., OR, USA) was used to calculate the mass concentrations for different species in NR-151 PM<sub>1</sub>. Default relative ionization efficiencies (RIE) were used for organics (1.4), nitrate (1.1) and chloride (1.4), respectively (Ng et al., 2011a). RIE of 5.8 for ammonium and 1.2 for 152 sulfate were determined by the IE calibrations of ammonium nitrate and ammonium 153 154 sulfate. Meanwhile, data were corrected for the particle collection efficiency (CE), due to 155 particle bounce on the vaporizer. CE can be affected by relative humidity, mass fraction of ammonium nitrate and particle acidity. In our measurement, the particles were generally 156 neutral and dried before sampling into the ACSM. CE was calculated as CE<sub>dry</sub> = max (0.45, 157 158 0.0833 + 0.9167 × ANMF), where ANMF refers to the ammonium nitrate fraction in NR-159 PM<sub>1</sub>(Middlebrook et al. 2012).

# 160 **2.2.2 OA source apportionment**

The receptor model PMF using a multilinear engine (ME-2) was used to identify and quantify the OA sources. PMF is a bilinear receptor model used to describe the variability of a multivariate dataset, X, as the linear combination of a set of constant factor profiles, F, and their corresponding time series G, as expressed in equation 1.

$$165 X = GF + E (1)$$

where X is the measured OA mass spectra consisting of *i* rows and *j* columns, and E is the model residuals. The PMF uses a least squares method to minimize the object function Q, defined as the sum of the squared residuals ( $e_{ij}$ ) weighted by their respective uncertainties ( $\sigma_{ij}$ ).

170 
$$Q = \sum_{i=1}^{m} \sum_{j=1}^{n} (e_{ij} / \sigma_{ij})^2$$
(2)

171 Unconstrained PMF analyses of OA data suffer from rotational ambiguity when sources 172 show similar profiles and temporal covariation (Canonaco et al., 2013; Huang et al., 2019). 173 However, by introducing *a priori* information as additional model input and constraining 174 one or more output factor profiles to a predetermined range, ME-2 can overcome such 175 difficulties and provide more environmentally meaningful solutions. When an element of 176 a factor profile ( $f_j$ , where j refers to the m/z) is constrained with a certain a value (a), the 177 following conditions need to be fulfilled:

178 
$$f_{j,solution} = f_j \pm a \times f_j$$
 (3)

The *a* value can vary between 0 and 1, which is the extent to which the output profiles can vary from the model inputs. The data analysis were conducted using the source finder (SoFi, Canonaco et al., 2013) tool version 4.9 for Igor Pro. Due to rotational ambiguity, there was no mathematically unique solution. Therefore, criteria including chemical fingerprint of the factor profiles, correlations with external tracers, and diurnal cycles were used for the factor identification and interpretation (Ulbrich et al., 2009; Huang et al., 2014, Elser et al., 2016).

## 186 **2.2.3 Aerosol liquid water content**

187NR-PM1 inorganic species, NH3 concentrations and meteorological parameters including188temperature and RH were used to calculate the aerosol liquid water content189inorganic species189(ALWCi) based on the ISORROPIA-II model (Fountoukis and Nenes,1902007). Here we ran the ISORROPIA-II in "forward" mode and the particles were assumed191to be deliquescent, i.e., in metastable mode (Hennigan et al., 2015). The thermodynamic192equilibrium of the NH4+-SO42--NO3--Cl--H2O system was then modeled and ALWCi was193calculated.

Meanwhile, the contribution of organics to ALWC (ALWC<sub>0</sub>) was also calculated using the
 following equation (Guo et al., 2015; Cheng et al., 2016):

$$W_{\rm org} = \frac{\rm OM}{\rho_{\rm org}}.\rho_{\rm w}.\frac{\kappa_{\rm org}}{(100\%/\rm RH-1)}$$

197 where OM is the mass concentration of organics,  $\rho_w$  is the density of water and  $\rho_{org}$  is the 198 density of organics ( $\rho_{org} = 1.4 \times 10^3$  kg m<sup>-3</sup>, Cerully et al., 2015).  $\kappa_{org}$  is the hygroscopicity 199 parameter of organic aerosol composition. We adopted a  $\kappa_{org}$  value of 0.06 based on 200 previous cloud condensation nuclei measurements in Beijing (Gunthe et al., 2011).

### 201 3 Results and discussion

# 202 **3.1 Temporal variations and mass fractions of PM<sub>1</sub> species**

203 Fig. 1 shows the time series of mass concentrations of OA, SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup>, Cl<sup>-</sup>, and BC, as 204 well as the meteorological parameters. The average mass concentration of  $PM_1$  during the 205 entire measurement period was 73.8 µg m<sup>-3</sup>, similar to those observed in Beijing in winter 206 2011 (66.8 μg m<sup>-3</sup>, Sun et al., 2013) and winter 2013 (64.0 μg m<sup>-3</sup>, Sun et al., 2016). The 207 lowest daily average concentration was 5.2 µg m<sup>-3</sup> on 31 December, while the highest was 210.1  $\mu$ g m<sup>-3</sup> on 15 January, with a difference of a factor of ~40. OA (52%) was the most 208 209 abundant component of  $PM_1$ , irrespective of the meteorological conditions, followed by 210 nitrate (14%) and sulfate (11%). The weather conditions during the measurement period 211 were characterized by drastic changes in wind speed, wind direction, RH and temperature, 212 providing a unique setting to investigate the influence of meteorological conditions on PM 213 species. The clean and pollution episodes occurred alternately during the measurement 214 period, and the PM1 concentration was usually lower than 20 µg m<sup>-3</sup> during clean episodes 215 and higher than 100 µg m<sup>-3</sup> during pollution episodes. As such, the measurements can be 216 divided into the clean period ( $PM_1 < 20 \ \mu g \ m^{-3}$ ) and the pollution period ( $PM_1 > 100 \ \mu g \ m^{-3}$ ) 217 <sup>3</sup>). South/southeasterly wind directions with low speed (average,  $0.9 - 1.0 \text{ m s}^{-1}$ ) were 218 typical for the pollution period, while north/northwesterly with high speed (average, 2.5 219 m s<sup>-1</sup>) for the clean period (Table Fig. 1).

220 During the polluted period, RH varied from 15% to 95% with an average value of 46% and 221 a median value of 43%. To investigate the effects of RH on PM pollution formation, we further divided the pollution period into two categories, the low-RH pollution days (RH 222 223 <50%) and the high-RH pollution days (RH >50%). The diurnal variations of mass 224 concentrations and fractions of different chemical species during clean days, low-RH 225 pollution days and high-RH pollution days are shown in Fig. 2. The mass fractional 226 variations were flatter during low-RH and high-RH pollution days than during clean days, 227 likely due to the accumulation of pollutants during stagnant weather conditions in pollution days. During clean days, secondary inorganic aerosol showed generally 228 229 increasing trends from 06:00 to 20:00 local time (LT), despite the development of the 230 boundary layer height during the day. The growth rate of nitrate mass (0.21  $\mu$ g m<sup>-3</sup> h<sup>-1</sup>) 231 was higher than that of sulfate  $(0.04 \,\mu g \, m^{-3} \, h^{-1})$  and ammonium  $(0.10 \,\mu g \, m^{-3} \, h^{-1})$ , indicating 232 that formation of nitrate was perhaps faster than that of sulfate and ammonium during 233 clean days. During low-RH pollution days, nitrate increased from 06:00 to 20:00 LT, with a growth rate of 0.40  $\mu$ g m<sup>-3</sup> h<sup>-1</sup>, which was two times higher than that during clean days. 234

On the contrary, sulfate concentrations increased from 06:00 to 10:00 LT, then started 235 decreasing and reached the minimum at 14:00 LT, possibly due to the increase of the 236 boundary layer height during the day, which outweighed the production of sulfate. 237 238 Associated with both sulfate and nitrate, ammonium showed a minor increase from 06:00 to 20:00 LT with a mass growth rate of 0.18  $\mu$ g m<sup>-3</sup> h<sup>-1</sup>. This phenomenon suggested that 239 240 the low-RH condition was favorable for nitrate formation but not for sulfate formation under polluted conditions. In contrast, obvious increases of secondary inorganic species 241 from 8:00 to 16:00 LT were observed during high-RH pollution days, with growth rates of 242 0.81  $\mu$ g m<sup>-3</sup> h<sup>-1</sup>, 0.55  $\mu$ g m<sup>-3</sup> h<sup>-1</sup> and 0.46  $\mu$ g m<sup>-3</sup> h<sup>-1</sup> for sulfate, nitrate and ammonium, 243 244 respectively. These mass growth rates increased correspondingly by about 20, 2.6 and 4.6 245 times compared to those during clean days. Note that nitrate growth rate in high-RH 246 pollution days (0.55  $\mu$ g m<sup>-3</sup>h<sup>-1</sup>) was still slightly higher than that in low-RH pollution days 247 (0.40  $\mu$ g m<sup>-3</sup> h<sup>-1</sup>), indicating that nitrate production is still efficient when RH is high, although not as much higher compared to sulfate. Measurements of sulfate oxygen 248 isotopes suggest that the largely enhanced formation of sulfate is associated with efficient 249 250 aqueous-phase reactions during high-RH pollution days (Shao et al., 2019). Note that the 251 comparison of growth rates was done under the assumption that chemical processes 252 were the main reason for mass growth, which might not be the case if other factors such 253 as planetary boundary layer height variations dominate. Yet comparison of growth rates of different species in the same time period would not be affected by these factors because 254 255 those species should share the same effects.

# 256 **3.2 Sources and diurnal variations of OA**

257 Source apportionment was performed on the OA data. Three to seven factors were 258 examined using an unconstrained PMF model, and the factors were qualitatively identified based on their mass spectral profiles and correlation with external data. We found that a 259 solution of five factors (i.e., HOA, COA, CCOA, BBOA, and OOA) best explains our data. For 260 261 the solutions with less than 5 factors, HOA appeared to be mixed with COA while CCOA mixed with BBOA (Fig. S1). However, when the number of factors was increased to 6, the 262 OOA factor split into two OOA factors of similar time series (Fig. S2), suggesting that 263 264 further separation of the factors does not improve the interpretation of the data.

265 Although five factors with different profiles and temporal variations were identified by 266 the unconstrained PMF model, the factor profiles and time series were suboptimal, specifically for HOA, COA, and BBOA. The diurnal pattern of HOA showed pronounced 267 peaks at cooking time, indicating its mixing with COA. The fractional contribution of m/z 268 60 ( $f_{60}$ , typically related to the fragmentation of anhydrous sugars) in HOA (0.008) was 269 higher than the average value reported from multiple ambient datasets (0.002, Ng et al., 270 2011). To reduce the mixing between factors, the reference HOA mass spectral profile, 271 272 characterized by a small  $f_{60}$  (Wang et al., 2017), and the BBOA mass spectral profile, 273 derived from Beijing wintertime measurements (Elser et al., 2016), were constrained 274 using ME-2. For the COA mass spectral profile that was derived from our unconstrained PMF analysis, a-value of 0 was used. Meanwhile, for HOA and BBOA, the a values were 275

varied systematically between 0 and 1 with an interval of 0.1 to explore the solution space. To assess the obtained solutions, we have set thresholds for the highest acceptable  $f_{60}$ value (0.006) for HOA and  $f_{57}$  value (0.042) for BBOA, based on mass spectra obtained at multiple sites (mean ± 2 $\sigma$ , Ng et al., 2011). Only solutions that conform to both criteria were selected and the final solution was the average of those selected reasonable solutions (Fig. S3).

The final OA factors resolved by ME-2 include four POA (i.e., HOA, COA, BBOA and CCOA), 282 283 and one SOA (i.e., OOA) factors, on average accounting for 14%, 14%, 10%, 32% and 31% 284 of OA mass concentration, respectively. The mass spectral profiles and time series of the 285 resolved factors are shown in Fig. 3a and b, respectively. The diurnal patterns of these 286 factors are presented in Fig. 4. The HOA spectrum is similar to those derived from other 287 studies in Beijing (Hu et al., 2016; Sun et al., 2014; 2016) and Pittsburgh (Ulbrich et al., 2016), and also resembles the source profile from diesel exhausts (Canagaratna et al., 288 2004). A strong correlation between the time series of HOA and BC was observed 289 290  $(R^2=0.84)$ . The diurnal cycle of HOA was similar to those observed in other studies in 291 Beijing (Sun et al., 2011; 2013; 2014), showing higher mass concentrations during the 292 night than during the day, due to enhanced traffic emissions from heavy duty vehicles and 293 diesel trucks that are allowed to enter the inner city during the night.

294 Similar to HOA, the mass spectrum of COA also displayed high signals in odd fragments, 295 while the m/z 55/57 ratio (1.45) and m/z 41/43 ratio (1.6) were significantly higher 296 compared to those of the HOA factor profile (m/z 55/57=0.65, m/z 41/43=0.88). The COA 297 profile is similar to those resolved in previous studies in Beijing (Elser et al., 2016; Sun et 298 al., 2016), Paris (Crippa et al., 2013) and Zurich (Dey et al., 2004). The  $R^2$  between COA and m/z 55 time series was 0.73. The diurnal cycle of COA showed two prominent peaks 299 during lunch (12:00-13:00 LT) and dinner (18:00-19:00 LT) times, and the peak in the 300 evening was more pronounced than that at noon, consistent with a previous study in 301 302 Beijing (Sun et al., 2016). Furthermore, the diurnal variation of COA was more obvious 303 with much clear noon and evening peaks during clean days than during low-RH and high-304 RH pollution days, likely because the stagnant meteorological conditions during pollution days facilitated the accumulation of pollutants and thus weakened the diurnal fluctuation. 305

306 The BBOA mass spectrum showed a similar pattern as that extracted from Crippa et al. 307 (2014), with pronounced peaks at m/z 60 and 73, two distinct markers of biomass burning emissions (Lanz et al., 2007). BBOA also showed similar time series with a high 308 signal at m/z 60 ( $R^2$ =0.74). The diurnal cycle of BBOA showed a slight increase during the 309 310 night (18:00-24:00 LT), corresponding to nighttime burning for residential heating in clean days, while this diurnal cycle became much flat during low-RH and high-RH 311 pollution days, likely due to the stagnant meteorological conditions during pollution days. 312 313 On average, BBOA contributed 10% of the total OA, much less than that of CCOA (32%), 314 consistent with previous results in Beijing (Elser et al., 2016).

315 The profile of CCOA showed a moderate correlation with that resolved in Beijing in winter

2014 (Elser et al., 2016). Similar to previous studies, signals related to unsaturated 316 hydrocarbons, especially those at m/z 77, 91 and 115, contributed significantly to the total 317 CCOA signal. In addition, there was a strong correlation between CCOA and Cl<sup>-</sup> ( $R^2$ =0.82), 318 which was considered as a marker mainly from coal combustion emissions. The mass 319 concentration and mass fraction of CCOA were both significantly higher at night than 320 those during the day, which was observed both in clean days and pollution days. The 321 diurnal pattern suggests much stronger emissions from coal combustion at night, a 322 323 situation further deteriorated by a shallower boundary layer at night.

One secondary OA factor, namely OOA, was also resolved, characterized by an important contribution at m/z 44. The profile of OOA is also similar to those resolved in Ng et al. (2011) and Sun et al. (2013). OOA is correlated well with nitrate ( $R^2$ =0.89).-, and the diurnal cycle of OOA shows an increase from about 6:00 to 20:00 LT, indicating the contribution from photochemical production and accumulation of OOA. Note that the growth rate of OOA during low-RH pollution days (1.0 µg m<sup>-3</sup> h<sup>-1</sup>) was higher than that during high-RH pollution days (0.40 µg m<sup>-3</sup> h<sup>-1</sup>) and clean days (0.35 µg m<sup>-3</sup> h<sup>-1</sup>) (Fig. 4).

### 331 **3.3 Chemically resolved PM pollution**

Fig. 5 shows the mass fraction of PM<sub>1</sub> and OA during clean, low-RH and high-RH pollution 332 periods. OA was the dominant component in PM<sub>1</sub>, with an average concentration 333 334 increasing from 10.9 µg m<sup>-3</sup> during clean periods to 56.7 µg m<sup>-3</sup> during high-RH pollution 335 periods and further to 67.7 µg m<sup>-3</sup> during low-RH pollution periods. The corresponding mass fraction of OA was 56%, 46%, and 54%, respectively. The decrease of OA mass 336 337 fraction during pollution periods can be attributed to the increased formation of sulfate and nitrate, as demonstrated in the above section. Specifically, nitrate increased from 11% 338 (2.2 µg m<sup>-3</sup>) during clean periods to 14% (17.2 µg m<sup>-3</sup>) during high-RH pollution periods 339 and to 15% (18.8 µg m<sup>-3</sup>) during low-RH pollution periods, while sulfate increased from 340 341 10% (2.0 µg m<sup>-3</sup>) during clean periods to 17% (20.9 µg m<sup>-3</sup>) during high-RH pollution periods but decreased back to as low as 7% (8.8  $\mu$ g m<sup>-3</sup>) during low-RH pollution periods. 342 343 The increased formation of nitrate from clean to pollution periods, especially during low-RH pollution periods, is likely due to enhanced photochemical production, as discussed in 344 345 Lu et al. (2019) which shows fast photochemistry during wintertime haze events in 346 Beijing. Specifically, the atmospheric oxidation proxy  $(O_x=O_3+NO_2)$  increased from 39.2 347 ppb during clean periods to 47.8 ppb during high-RH pollution periods, and up to as high as 59.8 ppb during low-RH pollution periods. Meanwhile, the precursor gas for nitrate, 348 349 NO<sub>2</sub>, increased accordingly from 16.7 ppb during clean periods to 64.342.2 ppb during high-RH pollution periods and to 103.055.4 ppb during low-RH pollution periods. The 350 351 averaged PM<sub>1</sub> concentrations during high-RH (123.2  $\mu$ g m<sup>-3</sup>) and low-RH (125.4  $\mu$ g m<sup>-3</sup>) pollution periods were very similar, but a distinct difference lies in the sulfate and nitrate 352 353 fractions in these two types of pollution periods. We observed similar a much larger 354 contributions from nitrate during low-RH pollution periods than during and high-RH 355 pollution periods, which may be due to enhanced photochemical formation and also 356 contributions of N<sub>2</sub>O<sub>5</sub> uptake, and while a much larger contribution from sulfate during high-RH pollution periods than during low-RH pollution periods because of enhancedformation from aqueous-phase processes.

In terms of OA sources, CCOA and OOA were the major sources irrespective of the PM<sub>1</sub> 359 360 level. The mass fraction of CCOA in OA increased from 25% (2.8 μg m<sup>-3</sup>) during clean periods to 31% (17.6 µg m<sup>-3</sup>) during high-RH pollution periods and to 35% (23.7 µg m<sup>-3</sup>) 361 during low-RH pollution periods, indicating the importance of residential coal 362 combustion emissions during haze pollution in wintertime Beijing (Elser et al., 2016; Li 363 364 et al., 2017). OOA also increased significantly during pollution periods, from 4.1  $\mu$ g m<sup>-3</sup> to 365  $\sim$ 20 µg m<sup>-3</sup>. It should be noted that the average OOA mass concentrations were rather 366 similar during high-RH (19.8 µg m<sup>-3</sup>) and low-RH (18.3 µg m<sup>-3</sup>) pollution periods. However, 367 the OOA mass fraction in OA during the high-RH pollution period (35%) is higher than 368 that during the low-RH pollution period (27%), indicating an additional contribution of OOA from e.g., aqueous-phase oxidations during high RH condition, as discussed below. 369 The mass fraction of HOA in OA increased from 8% (0.8 µg m<sup>-3</sup>) during clean days to 13% 370 371 (8.8 µg m<sup>-3</sup>) during low-RH pollution days and further to 16% (9.1 µg m<sup>-3</sup>) during high-RH pollution days, suggesting an increased contribution of HOA in pollution days. The mass 372 373 fraction of HOA is similar to those measured in wintertime Beijing in 2011(14%, Hu et al., 2016) and in 2013 (11%, Sun et al., 2016). In contrast, the mass concentrations of COA 374 during low-RH pollution days (8.8 µg m<sup>-3</sup>) and high-RH pollution days (6.8 µg m<sup>-3</sup>) were 375 376 higher than that during clean days (2.0 µg m<sup>-3</sup>), but the mass fraction of COA in OA during 377 high-RH pollution days (12%) and low-RH pollution days (13%) were lower than that 378 during clean days (20%). A similar decrease of HOA contribution and increase of COA 379 contribution during clean days were also observed by Sun et al. (2016) in wintertime 380 Beijing in 2011. The highest contribution of BBOA was observed during low-RH pollution 381 days with a mass fraction of 12% (8.1  $\mu$ g m<sup>-3</sup>). The BBOA concentration during high-RH pollution days (3.4  $\mu$ g m<sup>-3</sup>) was higher than that during clean days (1.0  $\mu$ g m<sup>-3</sup>), but the 382 383 mass fraction of BBOA in OA during high-RH pollution days (6%) was lower than that 384 during clean days (10%).

385 The chemical composition and sources of PM<sub>1</sub> under different meteorological conditions (e.g., wind direction, wind speed and RH) in the seven pollution episodes ( $PM_1 > 100 \mu g$ 386 m<sup>-3</sup>) and seven clean episodes (PM<sub>1</sub> <20  $\mu$ g m<sup>-3</sup>) are shown in Fig. S4. Note that these 387 388 episodes in total accounted for 91% of the entire measurement period. The pollution 389 episodes were found to be associated with the air masses from south/southwest, while clean episodes were associated with the air masses from north/northwest. Meanwhile, 390 the pollution episodes were generally associated with higher RH and lower wind speeds 391 392 when compared to the clean episodes. The wind speeds were approximately three times higher in clean episodes than those in pollution episodes. For example, the lowest 393 concentration of PM<sub>1</sub> was 6.7 µg m<sup>-3</sup> in C6 period, corresponding to the highest wind speed 394 395  $(4.0 \text{ m s}^{-1})$  and the lowest concentrations (< 20 ppb) of inorganic gaseous precursors (SO<sub>2</sub>, 396 NH<sub>3</sub>, and NO<sub>x</sub>), while the highest PM<sub>1</sub> concentration of 169.0  $\mu$ g m<sup>-3</sup> was found at P5, 397 corresponding to a much lower wind speed (<1.0 m s $^{-1}$ ). The mass concentrations of OA increased from ~4.1-9.4 µg m<sup>-3</sup> during clean episodes to ~44.7-85.7 µg m<sup>-3</sup> during 398

pollution episodes. However, the contributions of OA to PM<sub>1</sub> showed a decreasing trend 399 400 from 48-59% during clean episodes to 44-57% during pollution episodes, and the corresponding contributions of secondary inorganic species increased from 29-34% 401 (~2.2-5.5  $\mu$ g m<sup>-3</sup>) to 27-47% (~25.5-62.1  $\mu$ g m<sup>-3</sup>), indicating a notable production and 402 403 accumulation of secondary inorganic aerosol during haze pollution episodes. In contrast, 404 the mass concentration of OOA increased from  $\sim$ 1.4-3.9 µg m<sup>-3</sup> during clean episodes to ~10.0-27.6  $\mu$ g m<sup>-3</sup> during pollution episodes, while the contribution of OOA to OA 405 decreased from 33-64% during clean episodes to 20-52% during pollution episodes. The 406 407 corresponding contribution of POA sources increased from 35-67% ( $\sim$ 1.2-4.7 µg m<sup>-3</sup>) to 408 38-80% (~13.9-58.7 µg m<sup>-3</sup>), suggesting that in general the emission and accumulation of 409 POA sources played an important role during haze pollution in this measurement 410 campaign.

Comparing the pollution episodes with different RH conditions (see Fig. S4), the mass 411 fraction of sulfate was much higher during high-RH pollution episodes (P3, P6 and P7, 15-412 413 21%) than during low-RH pollution episodes (P1, P2, P4 and P5, 6-8%). OOA also showed 414 a much higher contribution to OA during high-RH pollution events (62% for P6 and 50%) 415 for P7) than during low-RH pollution events (P1, P2, P4 and P5, 20-31%). These variations suggest the potential importance of aqueous-phase reactions on the formation of sulfate 416 and OOA, as discussed above. Further comparison of high-RH and low-RH pollution 417 episodes with similar PM levels (e.g., P2 and P6 with  $PM_1$  concentration of 98.8  $\mu$ g m<sup>-3</sup> and 418 419 99.6µg m<sup>-3</sup>, respectively) shows that secondary inorganic aerosol dominated PM<sub>1</sub> at high-420 RH pollution episode. Similarly, as for the high-RH and low-RH pollution episodes with similar OA levels, for example, P6 (44.7µg m<sup>-3</sup>) and P7 (46.3 µg m<sup>-3</sup>), OOA dominated the 421 422 particulate pollution (62% of OA) at high-RH pollution events due to efficient formation 423 of SOA. On the contrary, POA had increased contributions to PM pollution at low RH and stagnant weather conditions (from 38% of OA at high-RH pollution to 50% of OA at low-424 425 RH pollution), consistent with previous studies in other Chinese cities (e.g., Wang et al., 426 2017; Huang et al., 2019). These results indicate that meteorological conditions have 427 important effects on the particulate pollution.

## 428 **3.4 Formation of secondary aerosol**

429 The relationship between  $SO_{4^{2-}}$  and  $NO_{3^{-}}$  is investigated to elucidate the formation 430 processes of these two typical secondary inorganic aerosol species. The correlation between  $SO_4^{2-}$  and  $NO_3^{-}$  was weak for the entire pollution period, because of the varied 431 relative contribution of different formation processes during different periods. However, 432 433 better correlations between SO<sub>4</sub><sup>2-</sup> and NO<sub>3</sub><sup>-</sup> were found with different slopes when the data 434 were divided into low-RH (RH <50%) and high-RH (RH >50%) pollution periods (Fig. 6). 435 During low-RH pollution periods,  $NO_{3^{-}}$  and  $SO_{4^{2^{-}}}$  showed a good correlation (R<sup>2</sup> = 0.75) 436 with a ratio of 2.1, indicating a similar photochemical production process. Meanwhile, the 437 high ratio between NO<sub>3</sub>- and SO<sub>4</sub>- suggest the nitrate production is more efficient than 438 that of sulfate during low-RH pollution period. However, during high-RH pollution periods, the ratio of  $NO_3$  to  $SO_4^{2-}$  decreased significantly to 0.40 with a lower correlation coefficient 439

440 ( $R^2 = 0.53$ ). The degraded temporal correlation between nitrate and sulfate suggest 441 different formation pathway of nitrate and sulfate during high RH pollution periods. 442 Aqueous-phase production of  $SO_4^{2-}$  become important during those periods. Consistently, 443 Fig. 7 shows that the sulfate oxidation ratio ( $SOR = [SO_4^{2-}]/([SO_4^{2-}] + [SO_2])$ ) increased 444 exponentially with the increase of ALWC at RH >50%.

445 A strong correlation of the mass concentrations between OOA and  $NO_3^-$  was observed with R<sup>2</sup> of 0.84 (Fig. 8a), possibly explained by the dominant contribution of photochemical 446 447 production for both OOA and NO<sub>3</sub><sup>-</sup>. <u>Meanwhile, the O<sub>x</sub> concentration during low-RH</u> 448 pollution days (59.8 ppb) was higher than that during high-RH pollution days (47.8 ppb) 449 and clean days (39.2 ppb). With the higher Ox concentration (as a surrogate of oxidant level) under low-RH conditions, the daytime formation of OOA was more efficient and the 450 451 growth rate was higher during those low-RH pollution days than those during high-RH 452 pollution days and clean days. When considering the RH effect (color coded in Fig. 8a), it is found that the data are scattered around the regression line with uniform slope when 453 454 RH <70% but concentrated in a small area above the regression line when RH >70%, suggesting that the OOA formation at RH >70% is probably promoted by aerosol water. 455 456 This is further supported by the linear increase of OOA with increasing SO<sub>4</sub><sup>2-</sup> when RH 457 >70%, while the relationship between OOA and  $SO_4^{2-}$  was very scattered when RH <70% 458 (Fig. 8b).

#### 459 4 Conclusion

We conducted online measurements of  $PM_1$  in urban Beijing from 29 December 2014 to 460 461 27 February 2015. The average mass concentration of PM<sub>1</sub> was 73.8  $\mu$ g m<sup>-3</sup> and OA was the most important component of  $PM_1$  (52%), followed by nitrate (14%) and sulfate (10%). 462 Source apportionment of OA resolved five factors including HOA, COA, BBOA, CCOA, and 463 464 00A, in which CCOA (32%) and 00A (32%) were the most important sources to 0A. The 465 mass proportion of CCOA in OA showed a significant increase from clean period (25%) to pollution periods (31-35%), highlighting the important role of coal burning in haze 466 467 formation in wintertime Beijing. The meteorological conditions (WD, WS, and RH) have a significant impact on the chemical composition and evolution of PM<sub>1</sub> species. Nitrate had 468 a higher contribution during low-RH pollution days, implying the photochemical 469 470 oxidation process of nitrate formation. In contrast, the mass fraction of sulfate to PM<sub>1</sub> was 471 much higher during high-RH pollution episodes compared to those during low-RH pollution episodes. The data also showed the exponential increase of sulfate oxidation 472 ratio (SOR) with ALWC at high RH conditions. Both are consistent with the impacts of 473 aqueous-phase reactions on the formation of sulfate. As for the OOA formation, the strong 474 475 correlation between OOA and  $NO_{3}$  may be explained by the dominant role of 476 photochemical production on both species; aqueous-phase processes may add an 477 additional contribution to OOA formation under high RH condition, as indicated by the 478 linear increase of OOA with increasing  $SO_4^{2-}$  when RH >70%. These results provide 479 insights into the relative importance of photochemical oxidation and aqueous-phase 480 processes for secondary aerosol formation during haze pollution, demonstrating the 481 <u>significance of meteorological conditions in the formation of secondary aerosol.</u>

*Data availability.* Raw data used in this study are archived at the Institute of Earth
Environment, Chinese Academy of Sciences, and are available on request by contacting
the corresponding author.

- 485 *Supplement.* The Supplement related to this article is available online at
- 486 *Competing interests.* The authors declare that they have no conflict of interest.

*Author contributions.* RJH designed the study. Data analysis and interpretation were made
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## 497 **References**

- Alfarra, M. R., Prévôt, A. S. H., Szidat, S., Sandradewi, J., Weimer, S., Lanz, V. A., Schreiber, D.,
  Mohr, M., and Baltensperger, U.: Identification of the mass spectral signature of organic
  aerosols from wood burning emissions, Environ. Sci. Technol., 41, 5770-5777, 2007.
- An, Z., Huang, R.-J., Zhang, R., Tie, X., Li, G., Cao, J., Zhou, W., Shi, Z., Han, Y., Gu, Z., and Ji, Y.:
  Severe haze in northern China: A synergy of anthropogenic emissions and atmospheric processes, Proc. Natl. Acad. Sci., 116(18), 8657–8666, 2019.
- Canagaratna, M. R., Jayne, J. T., Ghertner, D. A., Herndon, S., Shi, Q., Jimenez, J. L., Silva, P. J.,
  Williams, P., Lanni, T., Drewnick, F., Demerjian, K. L., Kolb, C. E., and Worsnop, D. R.:.
  Chase studies of particulate emissions from in-use New York City vehicles, Aerosol Sci.
  Tech., 38(6), 555-573, 2004.
- Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q., Onasch, T.
  B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R., Trimborn, A. M.,
  Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P., and Worsnop, D. R.: Chemical
  and microphysical characterization of ambient aerosols with the Aerodyne aerosol
  mass spectrometer, Mass Spectro. Rev., 26(2), 185–222, https://doi.org/10.100/mas.
  20115, 2007.
- Canonaco, F., Crippa, M., Slowik, J. G., Baltensperger, U., and Prévôt, A. S. H.: SoFi, an IGORbased interface for the efficient use of the generalized multilinear engine (ME-2) for
  the source apportionment: ME-2 application to aerosol mass spectrometer data,
  Atmos. Meas. Tech., 6, 3649–3661, https://doi.org/10.5194/amt-6-3649-2013, 2013.
- Canonaco, F., Slowik, J. G., Baltensperger, U., and Prévôt, A. S. H.: Seasonal differences in
  oxygenated organic aerosol composition: implications for emissions sources and
  factor analysis, Atmos. Chem. Phys., 15, 6993–7002, https://doi.org/10.5194/acp-156993-2015, 2015.
- 522 <u>Cerully, K. M., Bougiatioti, A., Hite Jr., J. R., Guo, H., Xu, L., Ng, N. L., Weber, R., and Nenes, A.:</u>
   523 <u>On the link between hygroscopicity, volatility, and oxidation state of ambient and</u>
   524 <u>water-soluble aerosols in the southeastern United States, Atmos. Chem. Phys., 15,</u>
   525 <u>8679–8694, https://doi.org/10.5194/acp-15-8679-2015, 2015.</u>
- 526 Cheng, Y. F., Zheng, G. J., Wei, C., Mu, Q., Zheng, B., Wang, Z. B., Gao, M., Zhang, Q., He, K. B.,
  527 Carmichael, G., Pöschl, U., and Su, H.: Reactive nitrogen chemistry in aerosol water as
  528 a source of sulfate during haze events in China, Sci. Adv., 2, e1601530,
  529 https://doi.org/10.1126/sciadv.1601530, 2016.
- Chow, J. C., Bachmann, J. D., Wierman, S. S.G., Mathai, C.V., Malm, W. C., White, W. H., Mueller,
  P. K., Kumar, N., and Watson, J. G.: Visibility: Science and Regulation, J. Air. Waste.
  Manage., 52 (9), 973-999, 2002.
- Crippa, M., Decarlo, P. F., Slowik, J. G., Mohr, C., Heringa, M. F., Chirico, R., Poulain, L., Freutel,
  F., Sciare, J., Cozic, J., Di Marco, C. F., Elsasser, M., Nicolas, J., Marchand, Nicolas, Abidi, E.,
  Wiedensohler, A., Drewnick, F., Schneider, J., Borrmann, S., Nemitz, E., Zimmermann, R.,
  Jaffrezo, J.-L., Prévôt, A. S. H., and Baltensperger U.: Wintertime aerosol chemical
  composition and source apportionment of the organic fraction in the metropolitan
  area of Paris, Atmos. Chem. Phys., 13, 961–981, https://doi.org/10.5194/acp-13-9612013, 2013.

- 540 DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C., Gonin,
  541 M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez, J. L.: Field542 deployable, high-resolution, time-of-flight aerosol mass spectrometer, Anal. Chem.,
  543 78(24), 8281–8289, https://doi.org/10.1021/ac061249n, 2006.
- DeCarlo, P. F., Ulbrich, I. M., Crounse, J., de Foy, B., Dunlea, E. J., Aiken, A. C., Knapp, D.,
  Weinheimer, A. J., Campos, T., Wennberg, P. O., and Jimenez, J. L.: Investigation of the
  sources and processing of organic aerosol over the Central Mexican Plateau from
  aircraft measurements during MILAGRO, Atmos. Chem. Phys., 10, 5257–5280,
  https://doi.org/10.5194/acp-10-5257-2010, 2010.
- Deng, X., Tie, X., Wu, D., Zhou, X. J., Bi, X. Y., Tan, H. B., Li, F., and Jaing, C. L.: Long-term trend
  of visibility and its characterizations in the Pearl River Delta (PRD) region, China,
  Atmos. Environ., 42(7), 1424-1435, 2008.
- Elser, M., Huang, R. J., Wolf, R., Slowik, J. G., Wang, Q., Canonaco, F., Li, G., Bozzetti, C.,
  Daellenbach, K. R., Huang, Y., Zhang, R., Li, Z., Cao, J., Baltensperger, U., El-Haddad, I.,
  and Prévôt, A. S. H.: New insights into PM2.5 chemical composition and sources in two
  major cities in China during extreme haze events using aerosol mass spectrometry,
  Atmos. Chem. Phys., 16, 3207–3225, https://doi.org/10.5194/acp-16-3207-2016,
  2016.
- Forster, P., Ramaswamy, V., and Artaxo, P.: Changes in atmospheric constituents and in
  radiative forcing, Cambridge University Press: Cambridge, United Kingdom, pp 129234, 2007.
- Goldstein, A. H., and Galbally, I. E.: Known and unexplored organic constituents in the
  earth's atmosphere, Environ. Sci. Technol., 41 (5), 1514-1521, 2007.
- Gunthe, S. S., Rose, D., Su, H., Garland, R. M., Achtert, P., Nowak, A., Wiedensohler, A., Kuwata,
   M., Takegawa, N., Kondo, Y., Hu, M., Shao, M., Zhu, T., Andreae, M. O., and Poschl, U.:Cloud
   condensation nuclei (CCN) from fresh and aged air pollution in the megacity region of
   Beijing, Atmos. Chem. Phys., 11, 11023–11039, 2011.
- Guo, H., Xu, L., Bougiatioti, A., Cerully, K.M., Capps, S.L., Hite, J.R., Jr, Carlton, A.G., Lee, S.H.,
   Bergin, M.H., Ng, N.L.: Fine-particle water and pH in the southeastern United States,
   Atmos. Chem. Phys., 15, 5211-5228, 2015.
- Hennigan, C. J., Izumi, J., Sullivan, A. P., Weber, R. J., and Nenes, A.: A critical evaluation of
  proxy methods used to estimate the acidity of atmospheric particles, Atmos. Chem.
  Phys., 15, 2775–2790, https://doi.org/10.5194/acp-15-2775-2015, 2015.
- Hagler, G. S. W., Bergin, M. H., Salmon, L. G., Yu, J. Z., Wan, E. C. H., Zheng, M., Zeng, L. M.,
  Kiang, C. S., Zhang, Y. H., Lau, A. K. H., and Schauer, J. J.: Source areas and chemical
  composition of fine particulate matter in the Pearl River Delta region of China, Atmos.
  Environ., 40 (20), 3802-3815, 2006.
- Han, S., Kondo, Y., Oshima, N., Takegawa, N., Miyazaki, Y., Hu, M., Lin, P., Deng, Z., Zhao, Y.,
  Sugimoto, N., and Wu, Y.: Temporal variations of elemental carbon in Beijing, J. Geophys.
  Res. Atmos., 114, 2009.
- He, L.-Y., Huang, X.-F., Xue, L., Hu, M., Lin, Y., Zheng, J., Zhang, R., and Zhang, Y.-H.: Submicron aerosol analysis and organic source apportionment in an urban atmosphere in Pearl
  River Delta of China using high-resolution aerosol mass spectrometry, J. Geophys. Res. Atmos., 116, D12, https://doi.org/10.1029/2010JD014566, 2011.

- Hien, P. D., Bac, V. T., and Thinh, N. T. H.: PMF receptor modelling of fine and coarse PM 10,
  in air masses governing monsoon conditions in Hanoi, northern Vietnam, Atmos.
  Environ., 38(2), 189-201, 2004.
- Hu, W. W., Hu, M., Yuan, B., Jimenez, J. L., Tang, Q., Peng, J. F., Hu, W., Shao, M., Wang, M.,
  Zheng, L. M., Wu, Y. S., Gong, Z. H., Huang, X. F., and He, L. Y.: Insights on organic aerosol
  aging and the influence of coal combustion at a regional receptor site of central eastern
  China, Atmos. Chem. Phys., 13, 10095–10112, https://doi.org/10.5194/acp-1310095-2013, 2013.
- Hu, W., Hu, M., Hu, W., Jimenez, J. L., Yuan, B., Chen, W., Wang, M., Wu, Y., Chen, C., Wang, Z.,
  Peng, J., Zeng, L., and Shao, M.: Chemical composition, sources, and aging process of
  submicron aerosols in Beijing: Contrast between summer and winter, J. Geophys. Res.
  Atmos., 121(4), 1955–1977, https://doi.org/10.1002/2015JD024020, 2016.
- Huang, R. J., Zhang, Y. L., Bozzetti, C., Ho, K. F., Cao, J. J., Han, Y. M., Daellenbach, K. R., Slowik,
  J. G., Platt, S. M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M.,
  Ciarelli, G., Piazzalunga, A., Schwikowski, M., Abbaszade, G., Schnelle-Kreis, J.,
  Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., Haddad, I.E., and Prevot, A.S.H.:
  High secondary aerosol contribution to particulate pollution during haze events in
  China, Nature, 514, 218–222, 2014.
- 602 Huang, R.-J., Wang, Y., Cao, J., Lin, C., Duan, J., Chen, Q., Li, Y., Gu, Y., Yan, J., Xu, W., Fröhlich, R., Canonaco, F., Bozzetti, C., Ovadnevaite, J., Ceburnis, D., Canagaratna, M. R., Jayne, J., 603 604 Worsnop, D. R., El-Haddad, I., Prévôt, A. S. H., and O'Dowd, C. D.: Primary emissions 605 versus secondary formation of fine particulate matter in the most polluted city 19, 606 in North China, Chem. Phys., 2283-2298, (Shijiazhuang) Atmos. https://doi.org/10.5194/acp-19-2283-2019, 2019. 607
- Huang, X.-F, Yu, J. Z., He, L.-Y., and Yuan, Z. B.: Water-soluble organic carbon and oxalate in
  aerosols at a coastal urban site in China: Size distribution characteristics, sources, and
  formation mechanisms, J. Geophys. Res. Atmos., 111(D22), 2006.
- Huang, X.-F., He, L.-Y., Hu, M., Canagaratna, M. R., Kroll, J. H., Ng, N. L., Zhang, Y.-H., Lin, Y.,
  Xue, L., Sun, T.-L., Liu, X.-G., Shao, M., Jayne, J. T., and Worsnop, D. R.: Characterization
  of submicron aerosols at a rural site in Pearl River Delta of China using an Aerodyne
  High-Resolution Aerosol Mass Spectrometer, Atmos. Chem. Phys., 11, 1865–1877,
  https://doi.org/10.5194/acp-11-1865-2011, 2011.
- Huang, X. F., He, L. Y., Xue, L., Sun, T. L., Zeng, L. W., Gong, Z. H., Hu, M., and Zhu, T.: Highly
  time-resolved chemical characterization of atmospheric fine particles during 2010
  Shanghai world expo, Atmos. Chem. Phys., 12, 4897–4907, https://doi.org/10.5194/
  acp-12-4897-2012, 2012.
- Huang, X.-F., Xue L., Tian, X.-D., Shao, W. W., Sun, T. L., Gong, Z. H., Ju, W. W., Jiang, B., Hu, M.,
  and He, L. Y.: Highly time-resolved carbonaceous aerosol characterization in Yangtze
  River Delta of China: Composition, mixing state and secondary formation, Atmos.
  Environ., 64, 200-207, 2013.
- Huffman, J. A., Docherty, K. S., Aiken, A. C., Cubison, M. J., Ulbrich, I. M., DeCarlo, P. F., Sueper,
  D., Jayne, J. T., Worsnop, D. R., Ziemann, P. J., and Jimenez, J. L.: Chemically-resolved
  aerosol volatility measurements from two megacity field studies, Atmos. Chem. Phys.,

- 627 9, 7161–7182, https://doi.org/10.5194/acp-9-7161-2009, 2009.
- Jenkin, M. E.: Investigation of an oxidant-based methodology for AOT40 exposure
  assessment in the UK, Atmos. Environ., 94, 332–340, 2014.
- Jimenez, J. L., Jayne, J. T., Shi, Q., Kolb, C. E., Worsnop, D. R., Yourshaw, I., Seinfeld, J. H., Flagan,
  R. C., Zhang, X., Smith, K. A., Morris, J. W., and Davidovits, P.: Ambient aerosol sampling
  with an Aerosol Mass Spectrometer, J. Geophys. Res.-Atmos., 108, 8425,
  doi:10:1029/2001JD001213, 2003.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., 634 635 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., 636 Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, 637 C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, J., Huffman, J. A., 638 639 Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., 640 Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. 641 R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., 642 Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of 643 644 organic aerosols in the atmosphere, Science, 326, 1525-1529, https://doi.org/10.1126/science.1180353, 2009. 645
- Kadowaki, S.: On the nature of atmospheric oxidation processes of sulfur dioxide to sulfate
  and of nitrogen dioxide to nitrate on the basis of diurnal variations of sulfate, nitrate,
  and other pollutants in an urban area, Environ. Sci. Technol., 20(12), 86-93, 1986.
- Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van
  Dingenen, R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski,
  Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K.,
  Vignati, E., Stephanou, E. G., and Wilson, J.: Organic aerosol and global climate
  modelling: a review, Atmos. Chem. Phys., 5, 1053–1123, https://doi.org/10.5194/acp5-1053-2005, 2005.
- Khoder, M. I.: Atmospheric conversion of sulfur dioxide to particulate sulfate and nitrogen
  dioxide to particulate nitrate and gaseous nitric acid in an urban area, Chemosphere,
  49(6), 675-84, 2002.
- Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C., and Prévôt, A. S. H.:
  Source apportionment of submicron organic aerosols at an urban site by factor
  analytical modelling of aerosol mass spectra, Atmos. Chem. Phys., 7, 1503–1522,
  https://doi.org/10.5194/acp-7-1503-2007, 2007.
- Lanz, V. A., Prévôt, A. S. H., Alfarra, M. R., Weimer, S., Mohr, C., DeCarlo, P. F., Gianini, M. F. D.,
  Hueglin, C., Schneider, J., Favez, O., D'Anna, B., George, C., and Baltensperger, U.:
  Characterization of aerosol chemical composition with aerosol mass spectrometry in
  Central Europe: an overview, Atmos. Chem. Phys., 10, 10453–10471,
  https://doi.org/10.5194/acp-10-10453-2010, 2010.
- Lee, B. P., Li, Y. J., Yu, J. Z., Louie, P. K. K., and Chan, C. K.: Characteristics of submicron particulate matter at the urban roadside in downtown Hong Kong Overview of 4
  months of continuous high-resolution aerosol mass spectrometer measurements, J.
  Geophys. Res.-Atmos., 120 (14), 7040-7058, 2015.

- Li, H., Zhang, Q., Zhang, Q., Chen, C., Wang, L., Wei, Z., Zhou, S., Parworth, C., Zheng, B.,
  Canonaco, F., Prévôt, A. S. H., Chen, P., Zhang, H., Wallington, T. J., and He, K.: Wintertime
  aerosol chemistry and haze evolution in an extremely polluted city of the North China
  Plain: significant contribution from coal and biomass combustion, Atmos. Chem. Phys.,
  17, 4751–4768, https://doi.org/10.5194/acp-17-4751-2017, 2017.
- Li, Y. J., Lee, B. P., Su, L., Fung, J. C. H., and Chan, C. K.: Seasonal characteristics of fine
  particulate matter (PM) based on high resolution time-of-flight aerosol mass
  spectrometric (HR-ToF AMS) measurements at the HKUST Supersite in Hong Kong,
  Atmos. Chem. Phys., 15, 37–53, https://doi.org/10.5194/acp-15-37-2015, 2015.
- Lu, K., Fuchs, H., Hofzumahaus, A., Tan, Z., Wang, H., Zhang, L., Schmitt, S. H., Rohrer, F.,
  Bohn, B., Broch, S., Dong, H., Gkatzelis, G. I., Hohaus, T., Holland, F., Li, X., Liu, Y., Ma, X.,
  Novelli, A., Schlag, P., Shao, M., Wu, Y., Wu, Z., Zeng, L., Hu, M., Kiendler-Scharr, A.,
  Wahner, A., and Zhang, Y.: Fast Photochemistry in Wintertime Haze: Consequences for
  Pollution Mitigation Strategies, Environ. Sci. Technol, 53(18), 10676-10684, 2019.
- Massoli, P., Fortner, E. C., Canagaratna, M. R., Williams, L. R., Zhang, Q., Sun, Y. L., Schwab, J. 685 J., Trimborn, A., Onasch, T. B., Demerjian, K. L., Kolb, C. E., Worsnop, D. R., and Jayne, J. 686 T.: Pollution Gradients and Chemical Characterization of Particulate Matter from 687 688 Vehicular Traffic Near Major Roadways: Results from the 2009 Queens College Air Aerosol Sci. Tech., 1201-1218, 689 Quality Study in NYC, 46, doi:10.1080/02786826.2012.701784, 2012.Matson, P., Lohse, K. A., and Hall, S. J.: The 690 691 globalization of nitrogen deposition: Consequences for terrestrial ecosystems, Ambio, 692 31 (2), 113-119,2002.
- Murphy, J. G., Day, D. A., and Cleary, P. A.: The weekend effect within and downwind of
  Sacramento Part 1: Observations of ozone, nitrogen oxides, and VOC reactivity,
  Atmos. Chem. Phys., 7(20), 5327-5339, 2007.
- Nilsson, P. T., Eriksson, A. C., Ludvigsson, L., Messing, M. E., Nordin, E. Z., Gudmundsson, A.,
  Meuller, B. O., Deppert, K., Fortner, E. C., Onasch, T. B., and Pagels, J. H.: In-situ
  characterization of metal nanoparticles and their organic coatings using laservaporization aerosol mass spectrometry, Nano Research, 8 (12), 3780-3795, 2015.
- 700 Ng, N. L., Herndon, S. C., Trimborn, A., Canagaratna, M. R., Croteau, P. L., Onasch, T. B., Sueoer, 701 D., Worsnop, D. R., Zhang, Q., Sun, Y. L., and Jayne, J. T.: An Aerosol Chemical Speciation Monitor (ACSM) for routine monitoring of the composition and mass concentrations 702 45 703 of ambient aerosol, Aerosol Sci. Technol., (7), 770-784, 704 https://doi.org/10.1080/02786826.2011.560211, 2011a.
- Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Zhang, Q., Ulbrich, M., and Worsnop, D. R.: Realtime methods for estimating organic component mass concentrations from aerosol
  mass spectrometer data, Environ. Sci. Technol., 45, 910–916, https://doi.org/10.1021
  /es102951k, 2011b.
- Paatero, P., and Tapper, U.: Positive matrix factorization: A non-negative factor model with
  optimal utilization of error estimates of data values, Environmetrics, 5 (2), 111-126,
  1994.
- Peng, R. D., Dominici, F., Pastor-Barriuso, R., Zeger, S. L., and Samet, J. M.: Seasonal analyses of air pollution and mortality in 100 US cities, Am. J. Epidemiol., 161 (6), 585594, 2005.Pope, C. A., Burnett, R. T., Thun, M. J., Calle, E. E., Krewski, D., Ito, K., and

- Thurston, G. D.: Lung cancer, cardiopulmonary mortality, and long-term exposure to
  fine particulate air pollution, J. Am. Med. Assoc., 287, 1132–1141, 2002.
- Pudasainee, D., Sapkota, B., Bhatnagar, A., Kim, S. H., and Seo, Y. C.: Influence of weekdays,
  weekends and bandhas on surface ozone in Kathmandu valley, Atmos. Res., 95(2–3),
  150-156, 2010.
- Schauer, J. J., Rogge, W. F., Hildemann, L. M., Mazurek, M. A., Cass, G. R., and Simoneit, B. R.:
  Source apportionment of airborne particulate matter using organic compounds as
  tracers, Atmos. Environ., 30 (22), 3837-3855, 1996.
- Schneider, J., Weimer, S., Drewnick, F., Borrmann, S., Helas, G., Gwaze, P., Schmid, O.,
  Andreae, M.O. and Kirchner, U.: Mass spectrometric analysis and aerodynamic
  properties of various types of combustion-related aerosol particles, Int. J. Mass.
  Spectrom., 258(1-3), 37-49, 2006.
- Seinfeld, J. H., Pandis, S. N., and Noone, K.: Atmospheric chemistry and physics: from air
  pollution to climate change, Physics Today, 51, 88, 1998.
- Shao, J., Chen, Q., Wang, Y., Lu, X., He, P., Sun, Y., Shah, V., Martin, R. V., Philip, S., Song, S.,
  Zhao, Y., Xie, Z., Zhang, L., and Alexander, B.: Heterogeneous sulfate aerosol formation
  mechanisms during wintertime Chinese haze events: air quality model assessment
  using observations of sulfate oxygen isotopes in Beijing, Atmos. Chem. Phys., 19, 61076123, https://doi.org/10.5194/acp-19-6107-2019, 2019.
- Song, S., Nenes, A., Gao, M., Zhang, Y., Liu, P., Shao, J., Ye, D., Xu, W., Lei, L., Sun, Y., Liu, B.,
  Wang, S., and McElroy, M.: Thermodynamic modeling suggests declines in water
  uptake and acidity of inorganic aerosols in Beijing winter haze events during
  2014/2015–2018/2019, Environ. Sci. & Tech. Let., 6(12), 752-760, 2019.
- Sun, C., Lee, B. P., Huang, D., Jie Li, Y., Schurman, M. I., Louie, P. K. K., Luk, C., and Chan, C. K.:
  Continuous measurements at the urban roadside in an Asian megacity by Aerosol
  Chemical Speciation Monitor (ACSM): particulate matter characteristics during fall
  and winter seasons in Hong Kong, Atmos. Chem. Phys., 16, 1713–1728,
  https://doi.org/10.5194/acp-16-1713-2016, 2016.
- Sun, Y. L., Zhang, Q., Schwab, J. J., Demerjian, K. L., Chen, W. N., Bae, M. S., Hung, H. M.,
  Hogrefe, O., Frank, B., Rattigan, O. V., and Lin, Y. C.: Characterization of the sources and
  processes of organic and inorganic aerosols in New York city with a high-resolution
  time-of-flight aerosol mass apectrometer, Atmos. Chem. Phys., 11, 1581–1602,
  doi:10.5194/acp-11-1581-2011, 2011.
- Sun, Y. L., Wang, Z., Dong, H., Yang, T., Li, J., Pan, X., Chen, P., and Jayne, J. T.: Characterization
  of summer organic and inorganic aerosols in Beijing, China with an Aerosol Chemical
  Speciation Monitor, Atmos. Environ., 51, 250–259,
  doi:10.1016/j.atmosenv.2012.01.013, 2012.
- Sun, Y. L., Wang, Z. F., Fu, P. Q., Yang, T., Jiang, Q., Dong, H. B., Li, J., and Jia, J. J.: Aerosol composition, sources and processes during wintertime in Beijing, China, Atmos. Chem.
  Phys., 13, 4577–4592, https://doi.org/10.5194/acp-13-4577-2013, 2013.
- Sun, Y. L., Wang, Z. F., Fu, P. Q., Jiang, Q., Yang, T., Li, J., and Ge, X. L.: The impact of relative
  humidity on aerosol composition and evolution processes during wintertime in
  Beijing, China, Atmos. Environ., 77, 927-934, 2013.
- Sun, Y., Jiang, Q., Wang, Z., Fu, P., Li, J., Yang, T., and Yin, Y.: Investigation of the sources and

- evolution processes of severe haze pollution in Beijing in January 2013, J. Geophys. Res.
  Atmos., 119, 4380–4398, https://doi.org/10.1002/2014JD021641, 2014.
- Sun, Y. L., Wang, Z. F., Du, W., Zhang, Q., Wang, Q. Q., Fu, P. Q., Pan, X., Li, J., Jayne, J., and
  Worsnop, D. R.: Long-term real-time measurements of aerosol particle composition in
  Beijing, China: seasonal variations, meteorological effects, and source analysis, Atmos.

764 Chem. Phys., 15, 10149–10165, https://doi.org/10.5194/acp-15-10149-2015, 2015.

- Sun, Y., Du, W., Fu, P., Wang, Q., Li, J., Ge, X., Zhang, Q., Zhu, C., Ren, L., Xu, W., Zhao, J., Han, T.,
  Worsnop, D. R., and Wang, Z.: Primary and secondary aerosols in Beijing in winter:
  sources, variations and processes, Atmos. Chem. Phys., 16, 8309–8329,
  https://doi.org/10.5194/acp-16-8309-2016, 2016.
- Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.:
  Interpretation of organic components from Positive Matrix Factorization of aerosol
  mass spectrometric data, Atmos. Chem. Phys., 9, 2891–2918, https://doi.org/10.5194
  /acp-9-2891-2009, 2009.
- Vecchi, R., Marcazzan, G., Valli, G., Ceriani, M., and Antoniazzi, C.: The role of atmospheric
  dispersion in the seasonal variation of PM1 and PM2.5 concentration and composition
  in the urban area of Milan (Italy), Atmos. Environ., 38 (27), 4437-4446, 2004.
- Wang, Y. C., Huang, R. J., Ni, H. Y., Chen, Y., Wang, Q. Y., Li, G. H., Tie, X. X., Shen, Z. X., Huang,
  Y., Liu, S. X., Dong, W. M., Xue, P., Fröhlich, R., Canonaco, F., Elser, M., Daellenbach, K.R.,
  Bozzetti, C., Haddad, EI., and Cao, J. J.: Chemical composition, sources and secondary
  processes of aerosols in Baoji city of northwest China, Atmos. Environ., 158, 128–137,
  https://doi.org/10.1016/j.atmosenv.2017.03.026, 2017.
- Wu, Y. Z., Ge, X. L., Wang, J. F., Shen, Y. F., Ye, Z. L., Ge, S., Wu, Y., Yu, H., and Chen, M. D.:
  Responses of secondary aerosols to relative humidity and photochemical activities in an industrialized environment during late winter, Atmos. Environ., 193, 66-78, 2018.
- Xu, W. Q., Han, T. T., Du, W., Wang, Q. Q., Chen, C., Zhao, J., Zhang, Y. J., Li, J., Fu, P. Q., Wang, Z.
  F., Worsnop, D. R., and Sun, Y. L.: Effects of Aqueous-Phase and Photochemical
  Processing on Secondary Organic Aerosol Formation and Evolution in Beijing, China,
  Environ. Sci. Technol., 51(2), 762–770, https://doi.org/10.1021/acs.est.6b04498,
  2017.
- Xu, W. Q., Sun, Y. L., Wang, Q. Q., Zhao, J., Wang, J. F., Ge, X. L., Xie, C. H., Zhou, W., Du, W., Li,
  J., Fu, P. Q., Wang, Z. F., Worsnop, D. R., and Coe, H.: Changes in aerosol chemistry from
  2014 to 2016 in winter in Beijing: insights from high resolution aerosol mass
  spectrometry, J. Geophys. Res. Atmos., 124(2), 1132–1147, 2018.
- Zhang, Q., Alfarra, M. R., Worsnop, D. R., Allan, J. D., Coe, H., Canagaratna, M. R., and Jimenez,
  J. L.: Deconvolution and quantification of hydrocarbon-like and oxygenated organic
  aerosols based on aerosol mass spectrometry, Environ. Sci. Technol., 39 (13), 49384952, 2005.
- Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I.,and Dzepina, K.:
  Ubiquity and dominance of oxygenated species in organic aerosols in
  anthropogenically-influenced Northern Hemisphere midlatitudes, Geophys. Res. Lett.,
  34 (13), 2007.
- Zhang, Q., Meng, J., Quan, J., Gao, Y., Zhao, D., Chen, P., and He, H.: Impact of aerosol
  composition on cloud condensation nuclei activity, Atmos. Chem. Phys., 12, 3783-

- 803 3790, https://doi.org/10.5194/acp-12-3783-2012, 2012.
- Zhang, Y., Schauer, J. J., Zhang, Y., Zeng, L., Wei, Y., Liu, Y., and Shao, M.: Characteristics of
  particulate carbon emissions from real-world Chinese coal combustion, Environ. Sci.
  Technol., 42 (14), 5068-5073, 2008.
- Zhang, Y., Tang, L., Yu, H., Wang, Z., Sun, Y., Qin, W., and Ge, S.: Chemical composition,
  sources and evolution processes of aerosol at an urban site in Yangtze River Delta,
  China during wintertime, Atmos. Environ., 123, 339-349, 2015a.
- Zhang, Y. J., Tang, L. L., Wang, Z., Yu, H. X., Sun, Y. L., Liu, D., and Zhou, H. C.: Insights into
  characteristics, sources, and evolution of submicron aerosols during harvest seasons
  in the Yangtze River delta region, China. Atmos. Chem. and Phys., 15(3), 1331-1349.
- Zhang, Y. M., Zhang, X. Y., Sun, J. Y., Lin, W. L., Gong, S. L., Shen, X. J., and Yang, S.:
  Characterization of new particle and secondary aerosol formation during summertime
  in Beijing, China, Tellus B., 63(3), 382-394, 2011.
- Zhang, Y. J., Tang, L. L., Sun, Y. L., Favez, O., Canonaco, F., Albinet, A., Couvidat, F., Liu, D. T.,
  Jayne, J. T., Wang, Z., Croteau, P. L., Canagaratna, M. R., Zhou, H. C., Prevot, A. S. H., and
  Worsnop, D.R.: Limited formation of isoprene epoxydiols-derived secondary organic
  aerosol under NOx rich environments in Eastern China, Geophys. Res. Lett., 44(4),
  2035 2043, https://doi.org/10.1002/2016GL072368, 2017.
- Zheng, B., Zhang, Q., Zhang, Y., He, K. B., Wang, K., Zheng, G. J., Duan, F. K., Ma, Y. L., and
  Kimoto, T.: Heterogeneous chemistry: a mechanism missing in current models to
  explain secondary inorganic aerosol formation during the January 2013 haze episode
  in North China, Atmos. Chem. Phys., 15, 2031–2049, https://doi.org/10.5194/acp-152031-2015, 2015.
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Species	Clean	High-RH pollution	Low-RH pollution
PM <sub>1</sub> (μg m <sup>-3</sup> )	19.5	123.2	125.4
Org (µg m-3)	10.9 (56%)	56.7 (46%)	67.7 (54%)
SO <sub>4</sub> <sup>2-</sup> (μg m <sup>-3</sup> )	2.0 (10%)	20.9 (17%)	8.8 (7%)
NO <sub>3</sub> - (μg m <sup>-3</sup> )	2.2 (11%)	17.2 (14%)	18.8 (15%)
NH4+ (μg m <sup>-3</sup> )	1.8 (9%)	12.3 (10%)	11.3 (9%)
Cl- (µg m-3)	1 <u>.0</u> (5%)	7.4 (6%)	8.8 (7%)
BC (μg m <sup>-3</sup> )	1.7 (9%)	8.6 (7%)	10.0 (8%)
HOA (µg m <sup>-3</sup> )	0.8 (8%)	9.1 (16%)	8.8 (13%)
COA (µg m <sup>-3</sup> )	2 <u>.0</u> (20%)	6.8(12%)	8.8 (13%)
BBOA (µg m <sup>-3</sup> )	1 <u>.0</u> (10%)	3.4 (6%)	8.1 (12%)
CCOA (µg m-3)	2.8 (25%)	17.6 (31%)	23.7 (35%)
00A (μg m <sup>-3</sup> )	4.1 (37%)	19.8 (35%)	18.3 (27%)
O <sub>x</sub> (ppb)	39.2	47.8	59.8
NO <sub>2</sub> (ppb)	16.7	<u>64.342.2</u>	<del>103.0<u>55.4</u></del>
RH (%)	25 <u>.0</u>	60 <u>.0</u>	31 <u>.0</u>
WS (m s <sup>-1</sup> )	2.5	1 <u>.0</u>	0.9
Vis (Km)	15.7	6.5	6.7

Table1 Summary of the PM<sub>1</sub> composition, OA sources and meteorological conditions
 during different pollution periods.



**Figure 1.** Time series of (a) wind speed (WS) and wind direction (WD), (b) Temperature (Temp) and relative humidity (RH), (c) visibility and  $PM_1$ , (d) NR-PM<sub>1</sub> species (i.e., OA, SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup>, Cl<sup>-</sup> and BC; note that OA is halved clarity), (e) OA factors (i.e., HOA, COA, BBOA, CCOA and OOA), and (f) relative contribution of PM<sub>1</sub> species.



Figure 2. The diurnal variations of mass concentrations and relative contributions of PM<sub>1</sub>
components during clean days (a, d), low-RH pollution days (b, e) and high-RH pollution
days (c, f).



Figure 3. The mass spectra(a) and time series(b) of OA factors (HOA, COA, BBOA, CCOA, and OOA).







Figure 4. The diurnal variations of mass concentrations and relative contributions of OA
factors during clean days (a, d), low-RH pollution days (b, e) and high-RH pollution days
(c, f).



Figure 5. PM<sub>1</sub> chemical composition (a) and OA source composition (b) pie chart as well
as the mass concentrations of PM<sub>1</sub> species(c) and OA sources(d) during clean, High-RH
pollution and Low-RH pollution periods.



**Figure 6.** The relationship between  $SO_{4^{2-}}$  and  $NO_{3^{-}}$  during low-RH (RH <50%) and high-RH (RH >50%) pollution episodes.





Figure 8. Scatter plot between the mass concentration of OOA and NO<sub>3</sub><sup>-</sup> (colored by RH)
(a), and scatter plot between the mass concentration of OOA and SO<sub>4</sub><sup>2-</sup> (colored by RH)
(b).