

Measurement report: Leaf-scale gas exchange of atmospheric reactive trace species (NO₂, NO, O₃) at a northern hardwood forest in Michigan

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Abstract. During the Program for Research on Oxidants: PHotochemistry, Emissions, and Transport (PROPHET) campaign from July 21 to August 3, 2016, field experiments of leaf-level trace gas exchange of nitric oxide (NO), nitrogen dioxide (NO₂), and ozone (O₃) were conducted for the first time on the native American tree species *Pinus strobus* (eastern white pine), *Acer rubrum* (red maple), *Populus grandidentata* (bigtooth aspen), and *Quercus rubra* (red oak) in a temperate hardwood forest in Michigan, USA. We measured the leaf-level trace gas exchange rates and investigated the existence of an NO₂ compensation point of ~~1 ppb~~, hypothesized based on a comparison of a previously observed average diurnal cycle of NO_x (NO₂ + NO) concentrations with that simulated using a multi-layer canopy exchange model. Known amounts of trace gases were introduced into a tree branch enclosure and a paired blank reference enclosure. The trace gas concentrations before and after the enclosures were measured, as well as the enclosed leaf area (single-sided) and gas flow rate to obtain the trace gas fluxes with respect to leaf surface. There was no detectable NO uptake for all tree types. The foliar NO₂ and O₃ uptake largely followed a diurnal cycle, correlating with that of the leaf stomatal conductance. NO₂ and O₃ fluxes were driven by their concentration gradient from ambient to leaf internal space. The NO₂ loss rate at the leaf surface, equivalently, the foliar NO₂ deposition velocity toward the leaf surface, ranged ~~from~~ 0–3.6 mm s⁻¹ for bigtooth aspen, and 0–0.76 mm s⁻¹ for red oak, both of which are ~90% of the expected values based on the stomatal conductance of water. The deposition velocityies for red maple and white pine ranged ~~from~~ 0.3–1.6 mm s⁻¹ and ~~from~~ 0.01–1.1 mm s⁻¹, respectively, and were lower than predicted from the stomatal conductance, implying a mesophyll resistance to the uptake. Additionally, for white pine, the extrapolated velocity at zero stomatal conductance was 0.4 ± 0.08 mm s⁻¹, indicating a non-stomatal uptake pathway. The NO₂ compensation point was ≤60 ppt for all four tree species and indistinguishable from zero at the 95% confidence level. This agrees with recent reports for several European and California tree species but contradicts some earlier experimental results where the compensation points were found to be on the order of 1 ppb or higher. Given that the sampled tree types represent 80–90% of the total leaf area at this site, these results negate the previously hypothesized important role of a leaf-scale NO₂ compensation point. Consequently, to reconcile these findings, further detailed comparisons between the observed and the simulated in- and above-canopy NO_x concentrations, and the leaf- and canopy-scale NO_x fluxes, using the multi-layer canopy exchange model

with consideration of the leaf-scale NO_x deposition velocities as well as stomatal conductances reported here, are recommended.

1 Introduction

The reactive nitrogen species nitric oxide (NO) and nitrogen dioxide (NO_2) are key components in tropospheric oxidation chemistry, affecting air quality by triggering the production of ground-level ozone, secondary organic aerosol, and acid rain. Forests cover 27% of the world's land surface, and 34% of the land area of the United States (FAO, 2016), and are an important land cover type in the continental cycling of NO and NO_2 (collectively termed NO_x). In remote and relatively unpolluted forests, the main source of NO is biogenic emission from soil microbial nitrification and denitrification processes. Once it escapes the soil, NO is transported through the canopy by turbulent mixing that is coupled to the atmosphere above the forest.

During this time, NO participates in chemical reactions with trace species present in ambient air, primarily with ozone to form NO_2 . This happens on a relatively short time scale of tens to a few hundred seconds. During daytime, additional reactions may further transform NO_2 to other oxidized nitrogen species, but on a longer time scale (Min et al., 2014). Physical loss pathways of NO_x within the canopy include dry deposition and leaf stomatal and cuticular uptake. The relative differences in the time scales of the turbulent mixing and the chemical and physical sink processes determine the amount of NO_x removed within the canopy, with the remaining NO_x being released into the boundary layer. The relative differences in the time scales of the turbulent mixing and the chemical and physical sink processes determine the amount of NO_x removed within the canopy. Any NO_x remaining in the ambient air is subject to be transported to the free troposphere above the forest.

The effect of leaf stomatal and cuticular uptake on the release of soil-emitted NO_x through forest canopy to the atmosphere is described using an empirical parameter, the canopy reduction factor (CRF), introduced by Yienger and Levy (1995) for application in large-scale atmospheric chemistry studies that generally rely on the so-called “big-leaf” approach to represent atmosphere-biosphere exchange without considering the process inhomogeneity of the loss processes within the canopy. Based on a parameterization using leaf area index and stomatal area index, it was estimated that 25–55% of soil-emitted NO_x is lost within forest canopies annually or seasonally depending on forest type. Those estimates of the effective release of soil NO_x were further corroborated in a study by Ganzeveld et al. (2002) using, instead of the big-leaf approach, a multi-layer canopy exchange model in a chemistry-climate model. Additionally, by including the influences of wind speed, turbulence, and canopy structure when calculating the CRF, Wang et al. (1998) estimated that up to 70% of NO_x was removed within the canopy in Amazon in April, agreeing with earlier results (Jacob and Wofsy, 1990). Accounting for both forests and other types of ecosystems, Wang et al. (1998) also estimated the global average canopy reduction at 20%, versus 50% by Yienger and Levy (1995). More recently, Delaria et al. (2018) investigated NO_x exchange with the leaves of *Quercus agrifolia* (California live oak) and obtained deposition velocities of NO_2 and NO under light and dark conditions. Implementing these results in a multi-layer single-column model, it was calculated that California oak woodland canopy removes 15-30% of soil-emitted NO_x , and other forests in California and Michigan, close to 60% (Delaria and Cohen, 2020). Implementing these results in a multi-layer

single-column model, it was calculated the California oak woodland canopy removes 15–30% of soil-emitted NO_x , a significant amount but lower than the previously estimated numbers (Delaria and Cohen, 2019).

Similarly, vegetation and plant surfaces also affect ozone levels through dry deposition (Clifton et al., 2019, 2020; Silva and Heald, 2018; Kavassalis and Murphy, 2017). In forested areas, ozone dry deposition occurs through leaf stomata as well as non-stomatal pathways including cuticular uptake, and wet or dry leaf surface reactions, while some O_3 is also removed by gas-phase chemical reactions e.g. with biogenic volatile organic compounds (BVOCs) and NO . Though these processes have been identified, the exact partitioning between the dry deposition pathways (and in-canopy chemical destruction) has not been unequivocally determined, hindering the ability to correctly assess ground-level ozone. Thus, Forest canopy plays a significant role in regulating the trace gas compositions in the atmosphere. Direct observations of NO_x exchange with a wide variety of plants and in various ecosystems are necessary to achieve a better understanding of their overall impacts, better understand ecosystem impacts on NO_x cycling globally.

In fact, there have been over a dozen field and laboratory studies aimed at understanding leaf-level NO_x uptake conducted from since the 1990s to 2009, but primarily on European tree species (Raivonen et al., 2009), and references therein). From direct measurements of foliar NO_x uptake, a reasonably detailed understanding of the gas exchange processes between NO_x and O_3 and plant leaves has been developed. Plants absorb NO_2 and O_3 mainly through leaf stomata, but also by leaf cuticular uptake (Chaparro-Suarez et al., 2011; Coe, 1995; Geßler et al., 2002; Rondón et al., 1993). The uptake efficiency varies across plants and is influenced by environmental conditions. Studies at leaf level and within leaves have found that after entering the stomata, NO_2 is metabolized through dissolution and enzyme-catalyzed reactions (Hu et al., 2014; Nussbaum et al., 1993; Vallano and Sparks, 2008; Weber et al., 1998). Unlike NO_2 and O_3 , foliar exchange of NO is small to insignificant (Hereid and Monson, 2001; Rondón et al., 1993), except for herbicide-treated soybeans (Klepper, 1979) and nutrient-fed sugar cane, sunflower, corn, spinach, and tobacco plants (Wildt et al., 1997), where NO emission was observed. Results from (Delaria et al., 2018) are consistent with these earlier findings.

In addition to NO_x and O_3 deposition fluxes, NO_2 compensation points have also been obtained by extrapolating the linear relationship between NO_2 flux and the ambient NO_2 concentration over the leaf surface (Raivonen et al., 2009; Slovik et al., 1996). The compensation point is the specific NO_2 ambient mole fraction or concentration at which NO_2 uptake by the plant leaves or NO_2 flux toward the leaf surface becomes zero. Reported values for this NO_2 compensation point ranged from 0.3 to over 3 ppb, depending on tree type and the conditions under which the measurements were made. The existence of such a point implies that when ambient NO_2 is below these thresholds, for example, in remote, unpolluted forest areas where it is usually less than 1 ppb, the soil-emitted NO_x would not be efficiently removed by the forest canopy as necessary for balancing the NO_x budget in the overlaying atmosphere above the forest. In fact, for those relatively clean conditions, the forest foliar would provide an additional atmospheric NO_x source.

This conundrum, discussed by Lerda et al. (2000), seemed to be resolved in the past decade when additional leaf-scale experiments were made using a new chemiluminescent NO_x detector equipped with a highly NO_2 specific blue light converter NO_x measurement instrument (Breuninger et al., 2012, 2013; Chaparro-Suarez et al., 2011). The improved NO_2 detection

specificity of the instrument prevented artifacts caused by augmentation of the NO₂ signal from other nitrogen compounds such as nitrous acid (HONO), nitric acid (HNO₃), and peroxyacyl nitrates (PANs). These artifacts may have caused an observed reduction of NO_x uptake that led to the conclusion of an (inferred) compensation point. The above work, on several native European trees, showed either a lower NO₂ compensation point than previously measured, at 0.05 to 0.65 ppb, or values not significantly different from zero at the 95% confidence interval, and do not support the possibility of a foliar NO_x source. However, when analyzing the observed NO_x and O₃ concentrations in a North American hardwood forest at the University of Michigan Biological Station (UMBS) research site using a multi-layer canopy exchange model, Seok et al. (2013) found that the best agreement between simulated and measured NO_x concentrations was obtained when a 1 ppb NO₂ compensation point was invoked. Further analysis to assess the sensitivity of the simulated NO_x mixing ratios to the representation of soil NO emission, leaf surface photolysis of nitrate, or advection was not able to reproduce the observations especially regarding the diurnal cycle of NO_x.

In order to verify these findings regarding the potential role of an NO₂ compensation point for the UMBS site, we conducted further field experiments on leaf-level gas exchange in summer 2016. This work is the first direct observation of folia gas exchange of NO_x and O₃ on mature trees growing naturally in a North American forest. To our knowledge, there has been one early study on young seedlings of several American tree species (Hanson et al., 1989), and one recent study on seedlings of California live oak (Delaria et al., 2018). In this work, we used a branch enclosure technique to measure mainly NO₂, as well as NO, and O₃ exchange rates at the leaf surface of four locally dominant tree species, *Pinus strobus* (eastern white pine), *Acer rubrum* (red maple), *Populus grandidentata* (bigtooth aspen), and *Quercus rubra* (red oak). Results obtained from these measurements provide information to reassess the possibility of a foliar NO₂ source and the role of the canopy on NO_x and O₃ cycling at this forest site. In this paper, we use both “uptake” and “foliar deposition” when describing trace gas exchange at the foliar level. Both terms refer to the process of trace gas loss upon contact with the leaf surface; but generally, the subject of “uptake” is the plant whereas the subject of “deposition” is the trace gas.

2 Experiment

2.1 Site description

The experiments were carried out at the Program for Research on Oxidants: PHotochemistry, Emissions, and Transport (PROPHET) research site at UMBS, which occupies about 10,000 acres on the northern tip of the Lower Michigan Peninsula (45.56° N, 84.71° W, Fig. 1). The area was heavily logged until the end of the 19th century. It also experienced several severe wildfires from 1880 to 1920. Natural reforestation started when the location was acquired for the research station in 1909. Today, bigtooth aspen, trembling aspen (*Populus tremuloides*), red maple, red oak, and white pine dominate within about a 1-km radius of the PROPHET site, whereas within a 60-m radius of the site, there are more white pine trees and almost no trembling aspen.

The northern part of the peninsula is fairly remote. The air is free from anthropogenic pollutants unless meteorological conditions result in the advection of air masses from surrounding major cities: to the southwest, Chicago, IL and Milwaukee, WI; to the southeast, Detroit, MI; and to the east, Toronto, ON. During the field experiments, about 35% of the time, air masses were coming from these directions. However, since the enclosures were purged with scrubbed ambient air (see below), the direct influence of pollutants on the enclosed plant material was minimal.

The enclosure measurements were carried out from July 20 to August 3, 2016. Sky conditions were sunny to mixed sun and clouds most of the time. The average ambient temperature measured in the canopy (~2 m ~~from the ground~~) was 24°C during daylight, and 18°C at night, with maximum and minimum temperatures of 31°C and 12°C, respectively. There were two main rainfall events, with the most recent being two days prior to the start of the enclosure experiments. The average soil temperature was near 19°C ~~throughout~~ the experiment period, and the soil moisture decreased gradually after the rainfall. These conditions are within the normal ranges for this site in July.

2.2 Methods

Branch enclosure experiments were conducted sequentially on branches of white pine, red maple, bigtooth aspen, and red oak. The estimated ages of the white pine, red maple, and red oak trees were about 15 to 20 years, and the bigtooth aspen, 5–10 years. All tree branches were selected based on their sun exposure, accessibility, and size. The height of the enclosed branches ranged from 3 to 10 m above the ground.

The enclosure system was composed of three parts: the enclosures, the airflow system, and the trace gas measurement instruments (Fig. 2). The enclosure, essentially a flow chamber, was constructed using a 61 x 91 cm ~~bag made of Tedlar[®] bag~~ (polyvinyl fluoride) (Jensen Inert Products, Florida, USA) with three factory-installed 0.95 cm diameter ports to attach tubing and sensor wires. The branch was carefully enclosed by the bag so that it was situated as close to the middle of the bag as possible. The open end of the bag was then closed around and tied onto the main stem of the branch, tight enough to secure the enclosure when it was inflated by the purge air, but also with enough leakage to allow air to escape during purging. Each branch enclosure was paired with an identical enclosure assembly without any plant material as the background reference to account for wall effects and other factors that may affect trace gas concentrations. The reference enclosure was placed adjacent to the branch enclosure but without obstructing the sunlight to the enclosed tree leaves.

Air delivery and air sample lines (~~polytetrafluoroethylene or PTFE~~), each about 30 m long, were connected to the enclosures and ~~to the instruments that were housed in an air-conditioned trailer parked nearby~~ at the site. Between the trailer and the enclosures, the air and sample lines were bundled together and sheathed inside black flexible insulation hoses linked together end to end. The hoses were wrapped in aluminum foil to keep the sample lines from absorbing heat from sunlight.

Ambient air from outside the trailer and scrubbed free of dust, O₃, and NO_x was used as the purge gas. An oil-free air compressor (Medo USA, now Nitto Kohki USA) was used to pull the ambient air through an organic vapor/acid gas respirator cartridge (Magid, Illinois, USA), which functioned as a dust filter. Downstream of the compressor, the air was further filtered

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160 by an ozone scrubber (Thermo Fisher Scientific), activated charcoal, and a NO_x scrubber (Purafil, Inc., Georgia, USA). ~~After the filters, P~~polytetrafluoroethylene (PTFE) tubing (12.7 mm OD and 9.52 mm OD) ~~after the filters~~ was used to carry the air to the enclosure chambers. The tubing was connected to the port on the Tedlar bag at the end near the tip of the enclosed branch, opposite the bag opening. Inside the bag connected to the same port was an air distributor made from a loop of tubing (9.52 mm OD) with pinholes (~1.5 mm diameter) about 1 cm apart drilled along its entire length. This allowed even distribution and mixing of the purge air insider the enclosure. The flow rate of the purge air was maintained at 37 L m⁻¹. The volume of the inflated enclosure was ~57 L, giving the air a residence time of ~1.5 min. ~~Additional resident time of sample air due to the sample line (30 m, 3.175 mm ID) was ~ 6 sec.~~

For the trace gas exchange experiments, known amounts of NO₂, NO, or O₃ were added into the ~~purge-scrubbed~~ air stream. NO₂ and NO were from compressed standard gas cylinders (Scott-Marrin, Inc., California, USA), and O₃ was made in situ using a mercury Pen-Ray lamp O₃ generator (UVP, ~~L-L-C~~, California, USA) and compressed zero air. A KOFLO® type mixer was placed just downstream of the trace gas inlet to ensure even mixing of the added trace component with the scrubbed air.

NO, NO₂, O₃, CO₂, and H₂O mixing ratios before and after the enclosures were measured, with the air sample selected using a set of solenoid valves. The concentrations of these gases were calculated using the ideal gas law and the measured air temperature. The ~~sampling~~ time for ~~the branch enclosure each sample~~ was five minutes ~~each~~, alternating between the enclosure inlet and outlet. The reference enclosure inlet and outlet were sampled once an hour. The environmental conditions were also recorded, including ambient and enclosure temperatures (S-THB, Onset Computer Corp., Massachusetts, USA), leaf temperatures (thermocouple wire sensors, Omega Engineering, Connecticut, USA), relative humidity (S-THB, Onset), leaf wetness (~~for qualitative assessment of leaf conditions only~~) (S-LWA, Onset), and photosynthetically active radiation (PAR) (S-LIA, Onset). Standard commercially available instruments were used for O₃ (Model 49i, Thermo Fisher Scientific, USA) and CO₂/H₂O (LiCor 840, Li-Cor Corp., Nebraska, USA). NO and NO₂ were measured using a home-built chemiluminescence detector that utilizes the light-emitting reaction of NO with O₃ (Ryerson et al., 2000).

The NO_x instrument was programmed to run on 5-minute cycles, each with a one-minute measurement of zero air (UHP, Airgas, USA), followed by a ~~2.5~~-min measurement of NO, and a ~~2.5~~-min measurement of NO₂. In NO₂ mode, NO₂ was first converted to NO, then measured the same as in the NO mode. ~~The conversion was done using A~~an LED UV light source (L11921-500, Hamamatsu Photonics) ~~was used to photolyze NO₂ for the conversion~~. The peak light emission of this LED was at ~~386~~5 ± 5 nm, matching the absorption peak of NO₂ and minimizing the interference from the unwanted photolysis of HONO. ~~The NO₂ to NO conversion efficiency was ~ 0.68. Because the ambient air was scrubbed to remove O₃ (and other trace gases) before entering the enclosures, the effect of ambient O₃ on NO_x measurements was negligible.~~ A high-concentration (1.5 ppm) NO standard dynamically diluted with ultra-high purity zero air (Airgas, USA) was used to calibrate the NO measurement. For the NO₂ calibration, NO in the same diluted standard was partially converted to NO₂ by adding a controlled amount of ozone (generated in situ using a Pen-Ray ozone generator and 99.98% oxygen). The instrument calibration runs were initiated automatically about every 7 hours during regular operation. The overall 1σ-precision for a 5-minute measurement cycle was ~2 ppt for NO and ~4 ppt for NO₂. The accuracy of the NO ~~and~~ /NO₂ measurements was ~30 ppt.

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Water vapor at the enclosure inlet and outlet was measured using a LiCor 840. The instrument was calibrated using a LiCor dew point generator. The ambient relative humidity results from the Onset sensors were compared with the data from a nearby AmeriFlux tower (within 100 m) (Vogel, 2016a), and the agreement was within 3%. It was noticed that on particularly hot and humid afternoons, there was condensation of water in the sample line leading to the instruments. The condensed water was removed promptly with gentle warming of the affected section of the sample line. The data recorded during these times were excluded.

After installation, each set of branch and reference enclosures was first purged with scrubbed ambient air at least overnight and through the early morning hours (up to ~10:00 local time) to allow the branches to acclimate, and also to reduce the amount of any possible surface-deposited photochemically labile compounds that might interfere with the measurements (Raivonen et al., 2006). The gas exchange experiments were then started and carried out for the following 2 to 4 days. A known amount of the trace gas was introduced into the purge air flow. This included zero concentration, i.e. purging with scrubbed air **only** between the trace gas additions. The maximum mixing ratios of the trace gases in the purge air were kept within the range of typically observed ambient measurements, i.e. NO₂ <1.2 ppb, NO<300 ppt, and O₃<60 ppb.

The enclosed leaves were harvested after the completion of the measurements and immediately placed in an oven to be air-dried at 65°C. The leaf area was then measured by forming a monolayer of the dried leaves on graph paper. The enclosed single-sided leaf area for white pine, red maple, bigtooth aspen, and red oak was 0.35, 0.26, 0.11, and 0.44 m², respectively.

Leaf-level uptake or emission of the trace gas leads to a trace gas concentration difference between the enclosure inlet and outlet. In the enclosure, the flux of the trace gas with respect to the leaf surface is:

$$F_x = \frac{Q}{A}(c_o - c_i), \quad (1)$$

where F_x is the flux of the trace gas x ; c_i and c_o are the trace gas concentration measured at the enclosure inlet and outlet, respectively. **A negative flux indicates trace gas loss in the enclosure due to uptake.** Q is the purge air flow rate. A is the one-sided area of the enclosed leaves, as the stomata, the part of leaf anatomy most relevant to gas exchange, generally are located on the underside of tree leaves (Kirkham, 2014). A resulting flux with a negative sign reflects the loss of the trace gas at the leaf surface, and a positive flux, emission from the foliage. All the trace gas concentration changes through the branch enclosure ($c_o - c_i$) were corrected against the background obtained from the reference enclosure before the fluxes were determined according to Eq. (1). The detection limit of flux, i.e. the minimum absolute value above which the flux is significantly non-zero ($p < 0.05$, or at the 95% confidence level), was determined using the flux data obtained during the scrubbed air purge at nighttime when no emission from the leaves was expected because **generally** the leaf stomata are closed **at night**. These detection limits (Table 1) reflect the measurement precision of the instruments, variations of the actual enclosure conditions over time, and fluctuation of the purge air flow rate. **Nighttime transpiration in trees and shrubs has been measured in prior work, with reports of nighttime transpiration rates ranging from 0 to as much as 25% of the daytime value (Dawson et al., 2007), suggesting that leaf stomata may remain open at night for some plants. However, this possibility did not affect the**

above results as there was no evidence of a consistent concentration difference above zero between the enclosure outlet and inlet measurements.

2.3 Tree branch samples

The representative tree species were determined based on the basal and leaf area coverage within the 60 m radius of the research site. Tree branches for the enclosure experiments were selected for their accessibility from the ground. Preferences were given to those with adequate sun exposure and to mature trees whenever possible. Enclosed branches of white pine (*Pinus strobus*) and red maple (*Acer rubrum*) were ~7 m above ground and from trees that were over 10 m in height. The enclosed branches of red oak (*Quercus rubra*, ~6 m) and bigtooth aspen (*Populus grandidentata*, ~4.5 m) were ~3 m above ground.

3. Results

We first examine the results obtained when only scrubbed air (without any addition of NO_x or O_3) flowed through the enclosures. The NO and NO_2 fluxes during the scrubbed air purge are shown in Fig. 3 along with the PAR and leaf temperature measured at the same time. Each data point represents a 5-minute measurement. The data points in gray are indistinguishable from zero within the 95% confidence interval based on the detection limits listed in Table 1. Those outside this confidence interval are marked by black symbols. It is expected that after the plant enclosures are conditioned with the hours-long scrubbed air purge, there will be no signal of NO_x at the enclosure outlet unless there is a source within the enclosure to supply a detectable amount of NO_x . Indeed, for most of the day, there was no detectable amount of NO_x (or O_3) at the enclosure outlet. However, besides a few scattered data points that are outside the confidence interval, there also appears to be some consistent positive fluxes of NO_2 lasting around 30 min or less occurring around noon or early afternoon. Because of the relatively short duration of this NO_2 emission during the brightest time of the day, it is unclear whether the flux was due to emission from leaves, or due to the photolysis of any oxidized nitrogen substrate remaining on the leaf surface even after the initial overnight and morning purging. The mixing ratios corresponding to the observed fluxes were less than 30 ppt in each of the enclosures. Below we present the results from each of the trace gas addition experiments.

3.1 NO_2

Nitrogen dioxide (NO_2) was introduced into the purge air at different concentrations between zero to 40 nmol m^{-3} (~1 ppb). Generally, when NO_2 was added to the enclosure, there was a negative flux, indicating ~~an~~ uptake of the trace gas by the plant material (Fig. 4, top panel). The magnitude of the flux ~~was~~ proportional to the input NO_2 concentration. In addition, when the input concentration was held constant ~~over the course of for~~ several hours or overnight, the flux had a diurnal pattern, e.g. bigtooth aspen on July 30 to July 31. It was lowest at night, increased through the morning hours and peaked around midday before diminishing again toward nighttime. This behavior strongly suggests that the NO_2 uptake by these trees is in large part

255 controlled by leaf stomatal aperture and, at the same time, driven by the NO₂ concentration gradient from the air around the
leaf surface to the leaf internal space.

260 There are a couple of factors that complicated the NO₂ gas exchange experiment. First, the NO₂ standard used for delivering
NO₂ to the enclosure contains about 5% NO that was unavoidably added to the enclosure. Secondly, when there was intense
direct sunlight, some NO₂ in the enclosure was photolyzed. While corrections for these interferences were done using the
measurements from the reference enclosure, it is difficult to completely remove the artifact caused by NO₂ photolysis. This is
because the sunlight exposure of the two enclosures, although situated side-by-side, was often uneven, and the measurements
of the enclosures were done not simultaneously but sequentially. This problem is particularly pronounced for clear sky
conditions with strong contrasts in sunlit and shaded conditions inside the canopy. The branch enclosure was always positioned
to get more sun exposure than the reference enclosure if choices needed to be made. Therefore, the branch enclosure likely
265 received more sunlight overall, even though it might be more shaded during some measurement cycles. Generally, for the
periods of strong sunlight, there is residual NO after the correction against the reference enclosure is made. If we assume all
this is due to an underestimation of NO₂ photolysis and make a further correction by combining the changes of NO and NO₂,
the data quality is not improved while more noise is introduced to the data. Because of this and because we are not absolutely
certain about all possible sources of NO_x from the branch enclosures, we prefer to adhere with the correction using only the
270 reference enclosure measurements and view the resulting NO₂ flux as an upper bound, with possibly as much as 20%
overestimation under direct sunlight conditions, which accounts for ~16% of all data during the NO₂ exchange experiments.
It is necessary to point out here that there were a couple of factors that complicated the NO₂-gas exchange experiment. First,
it was found that the NO₂-standard used for this experiment contained about 5% NO, which was unavoidably added to the
enclosure along with the NO₂. However, the mole fraction of this NO was at most ~50 ppt (~2 nmol m⁻³), an amount that did
275 not induce any significant NO flux in the enclosure as shown in the NO gas exchange experiment. Second, there was photolysis
of NO₂ in the enclosure when it was under relatively intense sunlight. In theory, this effect is removed by the background
subtraction using the results from the reference enclosure. However, in practice, although the branch enclosure and the
reference enclosure were set up side-by-side, the amount of sunlight experienced by each chamber often was different due to
the enclosure bag sizes and the need to always set the reference on the relatively shaded side of the accessible space to avoid
280 blocking the sunlight to the enclosed branch. This arrangement likely leads to an underestimation and insufficient correction
for the photolysis of NO₂ in the branch enclosure, and consequently, an overestimation of the NO₂ uptake flux. When
extrapolating the linear relation of NO₂ flux vs. NO₂ concentration to determine the compensation point (see below), this flux
overestimation may cause an overestimation of the inferred compensation point. It is possible to use the observed apparent NO
flux from the photolysis of NO₂ to correct this error, but, because there was also NO released from the enclosure when only
285 the scrubbed air was added, and it is not well understood and quantified, directly using the NO flux to correct for the photolysis
effect may introduce an even larger error. Without any correction, even in the worst case, the maximum amount of NO flux
observed was only about 20% of the total NO₂ flux at the same time, from both the photolysis of NO₂ and possible NO emission
from the enclosure.

The size of the stomatal aperture, regulated by the plant's need to optimize photosynthesis and simultaneously minimize water loss, can be gauged by stomatal conductance of water using Eq. (2) (Weber and Rennenberg, 1996):

$$g_{H_2O} = F_{H_2O} / (C_{H_2O_leaf} - C_{H_2O_enclosure}), \quad (2)$$

in which the flux of water (F_{H_2O}) due to plant transpiration is calculated by applying the measured water concentration difference at the inlet and outlet of the enclosure to Eq. (1). $C_{H_2O_leaf}$, the water concentration inside the leaf air space, is calculated using the measured temperature of the enclosed leaves, assuming the air in the leaf internal space is saturated with water vapor. $C_{H_2O_enclosure}$, the water concentration of the branch enclosure, is evaluated using the measured enclosure relative humidity and temperature data. The resulting stomatal conductance of water, g_{H_2O} , for the four enclosed branches is shown in Fig. 5a. The stomatal conductance has a clear diurnal pattern, mainly following the daily cycles of sunlight and photosynthesis. The magnitude varies from tree to tree. The conductance of the white pine and the red maple branches were similar, ranging from near zero at night to about 3 mm s⁻¹, while the conductance of red oak was 0 to ~1 mm s⁻¹, and the bigtooth aspen, 0 to ~6 mm s⁻¹. When the conditions are such that the difference between the leaf and air temperatures is small and the enclosure humidity is high, the difference between $C_{H_2O_leaf}$ and $C_{H_2O_enclosure}$ is also reduced, increasing the uncertainty in g_{H_2O} . In our measurements, this happened mostly from dawn to sunrise, accounting ~10% of the total data points, where the $(C_{H_2O_leaf} - C_{H_2O_enclosure})$ was within one standard deviation from zero.

Knowing the stomatal conductance of water, the expected rate of NO₂ deposition through the plant stomata can be calculated. Across the stomata, the deposition is a diffusion-controlled process (Weber and Rennenberg, 1996; Weber et al., 1998), where the expected rate is the product of the stomatal conductance of water multiplied by the square root of the ratio of the molecular weight of water to the molecular weight of NO₂:

$$g_x = g_{H_2O} \times \sqrt{\frac{MW_{H_2O}}{MW_x}}, \quad (3)$$

where g_x represents the expected stomatal uptake rate for the trace gas species x (here $x = \text{NO}_2$), and MW represents molecular weight.

From the measurements, the leaf-level NO₂ deposition velocity, v_{dNO_2} , is the NO₂ flux toward leaf surface (F_{NO_2}) normalized to the corresponding NO₂ concentration in the enclosure (c_{o,NO_2}):

$$v_{dNO_2} = F_{NO_2} / c_{o,NO_2}. \quad (4)$$

If NO₂ deposition is exclusively controlled by stomatal uptake, agreement between the measured deposition velocity and the calculated stomatal uptake rate is expected, i.e. $v_{dNO_2} = g_{NO_2}$. If not, additional factors, such as internal mesophyll resistance (Gut, 2002; Thoene et al., 1996), or leaf cuticular adsorption (Coe, 1995; Geßler et al., 2002; Rondón et al., 1993), may also play a role with the former decreases and the latter increases the overall foliar deposition velocity.

In Fig. 5b, the measured foliar deposition velocity of NO₂ is plotted together with the calculated stomatal uptake rate for comparison. The agreement is generally good for all experiments, suggesting that the foliar deposition of NO₂ for these tree

species is indeed closely related to stomatal aperture. This also suggests that the effects of internal mesophyll resistance and cuticular uptake of NO₂ are relatively minor. The strength of correlation between NO₂ deposition and stomatal conductance is evaluated using the Pearson correlation coefficient, ρ , which has a possible value between -1 to 1, with a value of 0 indicating no correlation and a value that is away from zero indicating increasing positive or negative correlation. In Fig. 6, the foliar NO₂ deposition velocity is plotted against the stomatal conductance for each tree. The correlation coefficient for bigtooth aspen is 0.96 and for red oak is 0.85, both showing a strong positive correlation between the deposition velocity and stomatal conductance. The correlation is also evident but relatively weaker for the white pine ($\rho = 0.73$) and the red maple ($\rho = 0.71$).

The relationship between the deposition velocity and stomatal conductance is also examined using linear regression analysis. If NO₂ deposition is entirely controlled by stomata, the deposition rate at zero conductance ($g_{H_2O} = 0$), when the stomata are closed, should be zero; and the slope of the deposition rate vs. stomatal conductance should be equal to $\sqrt{\frac{MW_{H_2O}}{MW_{NO_2}}}$, or 0.62 (recall Eq. (2)). This relationship is shown in Fig. 6 with the solid blue line. The best fit and the 95% confidence bounds are represented by the red solid and dashed lines. Also listed in the figure are the slope (m), the intercept (b), and the r^2 value of each fit. The linear relationship for bigtooth aspen appears to be the tightest, with over 90% of the data variation can be explained by the fit. The intercept is nearly zero, and the slope of 0.56 is close to 0.62, making it reasonable to conclude that for the bigtooth aspen, stomatal uptake dominates NO₂ loss at the leaf surface. A similar conclusion can be made for red oak, where r^2 is 0.72, the slope and the intercept are 0.54 and 0.03, respectively.

The red maple is different. The slope of NO₂ deposition rate to stomatal conductance is 0.25, far less than 0.62. The data also appear to have more scatter. On the time series plot (Fig. 5a), the stomatal conductance on the morning of July 25 (from 8:30 to 12:30) shows high variability that is not reflected by the NO₂ deposition rate at the same time. Possibly an unknown measurement issue for water concentration during this time [or sources of water exchange at leaf surface other than stomata](#) [\(more on this in the Discussion\)](#) led to the high variability. However, excluding this portion of the observations and using only the data obtained prior to this time window, from 13:00 on July 24 to 8:00 on July 25, resulted in a modestly improved linear fit with a slope still below 0.3. For white pine, the same slope is 0.40, also lower than the expected value of 0.62 based on stomatal-controlled diffusion. These lower than expected slopes imply there may exist mesophyll resistance to NO₂ uptake for these tree species. Such resistance to stomatal uptake of NO₂ has previously been observed on some trees such as European *Picea abies* (Norway spruce) seedlings (Thoene et al., 1996), and Amazonian *Laetia corymbulosa* and *Pouteria glomerata* (Gut, 2002). However, in a separate study of Norway spruce seedlings (Rondón and Granat, 1994), no evidence of internal resistance to NO₂ stomatal uptake was found. Researches on CO₂ diffusion and H₂O transport into leaf internal spaces have revealed that mesophyll resistance is subject to environmental perturbations, and the responses among and within species can vary (Xiao and Zhu, 2017). It is reasonable to assume that the mesophyll resistance to NO₂ uptake may also subject to environmental conditions, and systematic observations under different conditions are needed [for obtaining to obtain](#) more general conclusions.

The y-intercept of the fitted line accounts for any possible additional foliar deposition when the stomata are closed and consequently the stomatal conductance is zero. Of all four trees studies, only white pine has an intercept significantly larger than zero at $0.43 \pm 0.09 \text{ mm s}^{-1}$, indicating a possible role of wet leaf surfaces and/or cuticular uptake. More about this in the Discussion section below. For white pine, the intercept is 0.43 ± 0.09 , indicating a possible role of wet leaf surfaces and/or cuticular uptake. We first examined the possibility of wet leaf surfaces in the enclosure. We calculated the enclosure dew point using the temperature and relative humidity recordings and compared it to the measured leaf temperature. The leaf temperature was higher than the dew point at all times during the experiments, excluding the possibility of wet leaf surface. Thus, cuticular adsorption remains a possible explanation for this uptake while the stomata are closed. Previous researches on Norway spruce (Rondón et al., 1993, Gebler et al., 2002) also gave conflicting conclusions on the existence of cuticular uptake of NO_2 . As with the mesophyll resistance, further investigations involving experiments over a longer time span and under different environmental conditions are necessary to better understand this possible pathway and its dependencies.

3.1.1 Compensation point of NO_2

To determine at what concentration the NO_2 flux becomes zero, we plot the NO_2 flux vs the enclosure NO_2 concentration in Fig. 7. The stomatal conductance of each data point is represented by the color scale, with cool to warm colors corresponding to the stomatal conductance from low to high in each enclosure. As expected, the flux increases with increasing NO_2 concentration in the air surrounding the leaves; and at a given concentration, the flux increases with stomatal conductance. For each enclosure, we selected the data points taken when the stomatal conductance was at least 50–60% of its maximum measured during the experiments, indicated by the large, warm-colored symbols in Fig. 7. These data were then fit with linear regression for flux vs NO_2 concentration. The intercepts of the best-fit regression line and the zero-flux line, representing the compensation point, are listed in Table 2. For all four tree types within the range of stomatal conductance considered, the inferred NO_2 compensation point is well below 100 ppt, and not distinguishable from zero within measurement uncertainties.

3.2 NO

Nitric oxide (NO) was added to the purge air at concentrations up to $\sim 10 \text{ nmol m}^{-3}$ ($\sim 250 \text{ ppt}$) to the white pine, red maple, and bigtooth aspen enclosures (Fig. 4b). Red oak was not included in this experiment. No significant NO flux toward the leaf surface was observed. This agrees with observations made on Scots pine (Rondón et al., 1993), corn leaves (Hereid and Monson, 2001), and *Quercus agrifolia* (Delaria et al., 2018). On the contrary, for the white pine, a positive flux up to $2.7 \text{ pmol m}^{-2} \text{ s}^{-1}$ from the enclosure was measured when NO was added (Fig. 4b, White Pine), indicating emission from the enclosed plant material. This flux also appears to increase with the enclosure NO mixing ratio. Although the photolysis of surface deposited nitrogen oxides may cause such positive NO flux, during the scrubbed air purge prior to the addition of NO, there was no significant NO emission from the pine enclosure. That said, this experiment was done only once in a span of five hours from late morning to early afternoon. We cannot absolutely rule out possible interference from nitrogen containing chemical components in the system.

3.3 O₃

Up to 2.2 μmol m⁻³ (~55 ppb) of ozone (O₃) was introduced to the enclosures. As in the case of NO₂, there was an O₃ loss within the enclosure, and it increased with input O₃ concentration (Fig. 4c). Shown in Fig. 5c is the comparison of the measured foliar O₃ deposition velocity to the expected stomatal uptake rate calculated using leaf stomatal conductance and the square root of the ratio of the O₃ and H₂O molecular weight, $\sqrt{\frac{MW_{H_2O}}{MW_{O_3}}}$. For the red maple, bigtooth aspen, and red oak, these two values agree reasonably well, implying that foliar ozone loss is mainly through leaf stomatal and closely related to stomatal conductance. Correlation analyses were not performed here due to the limited number of data points.

For white pine, the measured leaf-level O₃ deposition velocity is significantly greater than the expected stomatal uptake rate by a factor of 2 or more. It is known that on average up to 60% of ozone deposition in vegetated areas is through non-stomatal pathways (Clifton et al., 2019). Within a branch enclosure, non-stomatal pathways can include deposition to wet leaf surfaces (Zhou et al., 2017; Altimir et al., 2004), cuticular uptake, chemical reactions at the leaf surface (Jud et al., 2016; Fares et al., 2010) and in the gas phase with biogenic organic compounds (BVOCs). Estimation of the possible contribution from gas-phase reactions with BVOCs was made as follows. The upper bounds of typical emission rates at 30 °C and PAR level at 1000 μmol m⁻² s⁻¹ for monoterpenes and other BVOCs (excluding isoprene) are 3 and 5 μg C g⁻¹ h⁻¹, respectively (Guenther et al., 1994). The speciation of major BVOCs emitted by white pine at UMBS is based on Kim et al. (2011), including α- and β-pinene, limonene, linalool, α-humulene, and β-caryophyllene. Using the rate constants of the BVOCs with ozone reactions (Burkholder et al., 2015), and the residence time of 1.5 min in the enclosure plus ~6 sec in the sample line before reaching the detector, the estimated ozone loss due to gas-phase chemical reactions was less than 1%. Even with optimal light and temperature conditions for BVOC emission, the estimated gas-phase chemical removal would only be on the order of a few percent.

For the white pine, the measured leaf-level O₃ deposition velocity is significantly greater than the expected stomatal uptake rate. During the experiment, the input O₃ mole fraction was ~45 ppb, and the outlet O₃ mole fraction from the enclosure was only ~25 ppb. We first examined the enclosure humidity and leaf wetness to look for evidence of water condensation on leaf surface. It is known that a 'wet skin' condition would enhance ozone deposition (Altimir et al., 2004; Zhou et al., 2017). Again, the comparison between leaf temperature and the dew point does not support the possibility of water condensing on the pine needles, as the leaf temperature was consistently higher than the dew point. Another known O₃ loss process is chemical removal by gas-phase reactions with biogenically emitted reactive volatile organic compounds (BVOCs). It could be at least partially responsible for the additional ozone loss besides leaf stomatal uptake because conifers have relatively high emission rates of monoterpenes and sesquiterpenes, both of which are more reactive towards O₃ compared to isoprene, the main component of BVOCs emitted by many deciduous trees. Among typical BVOCs, β-caryophyllene has the fastest reaction rate constant with O₃, on the order of 1×10⁻¹⁴ cm³ molecule⁻¹ s⁻¹ (Burkholder et al., 2015). The residence time of the gas in the enclosure is 1.5 minutes. Thus, at least 20 ppb of such highly reactive BVOCs would be needed inside the enclosure to account for the

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additional ozone loss, a condition that is hard to fulfill as it would require unusually high BVOC emission rates. O₃ foliar deposition velocities higher than predicted by stomatal conductance have also been reported for other conifers (Norway spruce, 2.5 mm s⁻¹ and Scots pine, 3.1 mm s⁻¹, (Rondón et al., 1993)). Because leaf level O₃ removal may involve pathways of stomatal uptake, leaf surface cuticular uptake, gas phase chemical reactions, as well as chemical reactions at leaf surface (Fares et al., 2010; Jud et al., 2016), the observed results here could be due to one dominant pathway or more probably a combination of these sinks, whose dependencies on environmental conditions remain to be investigated.

4 Discussion

4.1 Foliar trace gas exchange: NO₂ and O₃

For two weeks during the summer PROPHET2016 campaign, we examined the leaf-level NO₂, NO, and O₃ gas exchange of four different tree species. This work provided a first insight into the general characteristics of the gas exchange of these North American trees in their natural habitat. The trees used in the enclosure measurement represent 80% of the total leaf area within a 1000 m radius of the research site, and 90% within the 60 m radius. It is evident from the results that bidirectional foliar gas exchange depends on the trace gas in question and tree type, and is influenced by a combination of diverse and complex environmental conditions, similar to the findings from previous studies mainly on European tree species and on annual plants (grasses and crops). The foliar uptake rates of NO₂ and O₃ vary from tree to tree and even within the same tree. Leaf stomatal conductance of H₂O emerges as a strong indicator of the uptake efficiency. The foliar NO₂ deposition of bigtooth aspen and red oak is almost entirely controlled by stomatal aperture. For red maple and white pine, the correlation coefficient is over 0.7, even though the measured NO₂ foliar deposition velocity is 40–50% of the predicted stomatal uptake rate. Except for white pine, the O₃ foliar deposition velocities of red maple, bigtooth aspen, and red oak all the studied trees also covary with stomatal conductance (Fig. 5c). Generally, the leaf-level NO₂ and O₃ deposition velocity can largely be inferred from the stomatal conductance of water only, as also concluded from earlier studies of European tree species (Breuninger et al., 2013; Rondón and Granat, 1994).

Thus, the factors controlling leaf stomatal conductance would in turn greatly influence NO₂ and O₃ deposition to the forest canopy in a forested environment. These factors include PAR level, ambient temperature, moisture, and soil conditions, as well as ambient CO₂ (Jarvis, 1976). Further, the capability of the foliar uptake of trace gases would also depend on the intrinsic characteristics of leaf stomata, such as their size and density on the leaf surface, determined by plant species and stage of maturity, and factors such as growth history, leaf age, tree height and the vertical location of the leaf on the tree (Kirkham, 2014; Schäfer et al., 2000; Sparks et al., 2001). In this work, the stomatal conductance of the bigtooth aspen was 3–5 times higher than that of the other trees. Biological features may have contributed to this difference. Compared with the other three trees in this work, the aspen was younger and smaller. The enclosed branch was in the upper part of the crown containing developing new leaves. Past measurements, albeit on different species, have shown that for the same species under similar environmental conditions, leaves of young trees generally have higher stomatal conductance than old ones (Niinemets, 2002;

Hubbard et al., 1999; Fredericksen et al., 1995; Yoder et al., 1994). Another possible reason for the observed high g_{H_2O} , while direct evidence has yet to be found, is the number of stomata. Many trees have stomata on only the lower (abaxial) leaf surface; however, trees that belong to the genus *Populus*, which includes aspen, are an exception. They have stomata on both sides (amphistomatous), a feature that allows increased photosynthetic rate and fast growth (Kirkham, 2014). If the bigtooth aspen leaves are indeed amphistomatous, a relatively high g_{H_2O} can be expected. We compared the environmental conditions of the enclosures. The integrated PAR exposure levels were similar. The daily variation of the relative humidity in the bigtooth aspen enclosure was not significantly different from the others. In contrast, the average daily temperature was 19.2 °C, cooler than the temperatures (23.9 °C, 22.6 °C, and 21.6 °C) in the other enclosures, similar to the average ambient air temperature outside the enclosure during the same time, 19.1 °C, and 23.6 °C, 22.4 °C, and 21.3 °C. The combined conditions of moisture and temperature led to a relatively low vapor pressure deficit (VPD) in the aspen enclosure, 0.8 kPa, compared to 1.2 kPa (white pine), 1.0 kPa (red maple), and 1.4 kPa (red oak) in the others. Generally, VPD and g_{H_2O} are inversely correlated and a low VPD corresponds to a relatively high g_{H_2O} (Hubbard et al., 1999; Urban et al., 2017a, 2017b). However, because here we are comparing different tree species, we consider the observed results to stem from the combination of biological and environmental factors. Further examination of these factors is beyond the scope of this paper, nevertheless, it would be beneficial to take this temporal and spatial variability and inhomogeneity into account in model parameterizations of trace gas dynamics since plant stomata are the main conduit of NO_x and O_3 deposition over vegetation. To look for potential causes of this large difference, we examined the environmental conditions of the enclosures. The integrated PAR exposure levels were similar. The daily variation of the relative humidity in the bigtooth aspen enclosure was not significantly different from the others. In contrast, the average daily temperature was 19.2°C, cooler than the temperatures (23.9°C, 22.6°C, and 21.6°C) in the other enclosures. The effect of temperature on the stomatal conductance is generally coupled with the effect of water vapor pressure deficit (VDP), with the combined influence possibly leading to an increase in stomatal conductance if the VDP effect is not yet inducing stomatal closure for further temperature increases (Urban et al., 2017a, 2017b). A more likely explanation, without excluding the possible environmental influence, comes from tree species, and leaf position and age. Although many trees have stomata on only the lower leaf surface (abaxial surface), trees that belong to the genus *Populus*, which includes aspen, are an exception. Leaves of these trees have stomata on both sides, a feature that allows increased photosynthetic rate and fast growth (Kirkham, 2014). Additionally, the aspen sampled in this study was smaller and younger than the other three trees. The enclosed branch was close to the top of the tree, with relatively young and developing leaves. This type of leaves generally has relatively high stomatal density compared to more mature leaves located lower on a tree. Likely, the aspen leaves had more stomata per unit area than the other tree leaves sampled in this work, which led to higher stomatal conductance and the higher rates of NO_2 and O_3 uptake through stomata. In future, this possibility should be further verified by comparing the stomatal counts on the sampled leaves, but unfortunately was beyond the scope of this work. Given the large dependence of NO_2 and O_3 uptake on the leaf stomata properties, it is of interest to consider how the factors influencing stomatal conductivity,

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giving its variability and inhomogeneity, could be better incorporated into models from leaf to canopy scale to improve the canopy-wide understanding and description of NO_x and O_3 dynamics.

When extrapolated to zero stomatal conductance, the deposition velocity of NO_2 to white pine was 0.43 mm s^{-1} (Figure 6a), implying deposition unrelated to leaf stomata, possibly to wet leaf surfaces and/or to leaf cuticula. This observation does not exclude the possible existence of these pathways when the stomata are open. A deposition velocity higher than expected based on the stomatal conductance would result if there is significant non-stomatal deposition. On the other hand, mesophyll resistance renders a lower deposition velocity than the expected value. There is no mechanistic reason why the deposition velocity associated with either a non-stomatal pathway or mesophyll resistance should remain constant or vary linearly with stomatal conductance. The relationship of deposition velocity, $v_{d\text{NO}_2}$, and stomatal conductance, $g_{\text{H}_2\text{O}}$, would remain essentially linear as long as stomatal deposition dominates or the non-stomatal deposition term is constant while mesophyll resistance is small. However, if mesophyll resistance is significant, it would limit the increase of $v_{d\text{NO}_2}$ with stomatal conductance.

To assess the role of wet leaf surfaces to non-stomatal deposition, we calculated the white pine enclosure dew point using the temperature and relative humidity data and compared it to the measured leaf temperature. The leaf temperature was always higher than the dew point during the experiments, excluding the possibility of a wet leaf surface from the condensation of pure water. However, a microscopic water film may nevertheless form at a relative humidity as low as 50% if there are hygroscopic deposits on the leaf surface (Sun et al., 2016; Burkhardt and Hunsche, 2013; Burkhardt and Eiden, 1994). The microscopic water film could potentially modify gas exchange rates of water-soluble trace gases in the air. Data from this work do not contain information that can be used to delineate the possibilities of trace gas dissolution into microscopic water films or cuticular uptake. Further investigations with appropriately designed experiments, better measurement precisions, longer observation time, and under different environmental conditions are necessary to delineate the various possible deposition pathways and their dependencies.

To put our measurement results in perspective, we compare our measured daytime maximum foliar deposition velocity of NO_2 with the results from previous studies on American trees (Table 3). Although the development stages of the trees, PAR, humidity and temperature conditions are different, the results are comparable, ranging from 0.76 to 1.6 mm s^{-1} from this work and 0.4 – 1.8 mm s^{-1} from earlier work. We also compare our results to those of several native European trees - scots pine, evergreen oak, common oak, European beech, and silver birch, measured under the conditions of $\text{PAR} = 900 \mu\text{mol m}^{-2} \text{ s}^{-1}$, maximum temperature 27.7°C , and relative humidity 31.2–99.9% (Breuninger et al., 2013). The maximum NO_2 deposition rates were ~ 0.5 – 1 mm s^{-1} for all but the birch tree, which was $\sim 1.5 \text{ mm s}^{-1}$. These numbers are also fairly similar to those of pine, maple and oak reported here. What stands out but without a direct or closely related comparison is the high rate of trace gas uptake by the aspen leaves. Although the comparisons show reasonable agreement, it is evident that the NO_2 (and O_3) foliar uptake is highly variable depending on a myriad of conditions, both environmental and intrinsic to tree species and its developmental stage. Measurement results and comparisons from different studies are probably also sensitive to experimental

protocols and environmental conditions. These factors should be taken into consideration if more comparisons are to be made in the future work.

515 4.1.1 NO₂ compensation point

Measured fluxes of NO₂ toward the leaf surface while the stomatal conductance was at least 50% of the observed maximum value were used to assess the possible existence of an NO₂ compensation point. It would have been indicated by a zero or positive NO₂ flux at significantly non-zero NO₂ concentrations defined by the measurement system detection limit (Table 1). We found no such evidence of an NO₂ compensation point for all the tree species measured in this work. Indeed, this lack of evidence of a compensation point is also supported by the fact that no significant, sustained NO₂ emission was observed while the enclosures were purged with the scrubbed air only. For all four trees in this study, the compensation point, if it exists at all, would be well below 150 ppt. Thus, this finding does not support the existence of a 1 ppb NO₂ compensation point as suggested in the previously mentioned combined NO_x concentration measurement and canopy exchange model study (Seok et al., 2013) to reach the best agreement between the simulated and observed NO_x concentrations above and within the forest canopy at the UMBS site. We would like to point out that the NO₂ flux may approach zero even at higher NO₂ concentrations if the stomata are not adequately open and the stomatal conductance is lower than the values used above (Fig. 7). However, because here the NO₂ uptake is mainly through stomata, such zero flux at relatively high NO₂ mixing ratios is not indicative of a compensation point; rather, it is from the reduced eapability-capacity of absorbing NO₂ under the-reduced stomatal conductance. Our result agrees with recent reports on several other tree species that an NO₂ compensation point is not observed above the detection limit of the measurement using improved NO₂-specific instruments with minimal interference from other nitrogen compounds (Breuninger et al., 2013; Chaparro-Suarez et al., 2011).

4.2 NO

There was no significant leaf-level deposition of NO for all the tree species studied here. Instead, relatively small NO emissions were detected from white pine when up to ~250 ppt NO was added to the enclosure. Delaria et al (2018) reached the same conclusion from their study on Quercus agrifolia. We searched for possible errors that might have led to this observation-the results but could not find an obvious explanation. Certainly, additional measurements are necessary to verify this observation. Using the leaf area index of white pine at UMBS, 0.11 m²/m² (Vogel, 2016b) and the maximum measured flux, 2.7 pmol m⁻² s⁻¹, we estimated the potential canopy-wide NO flux from this emission to be 0.3 pmol m⁻² s⁻¹, less than 10% of the reported minimum soil NO emission flux of 4–10 pmol m⁻² s⁻¹ at UMBS (Nave et al., 2011).

540 Although this observation seems counterintuitive, in previous publications, emission of NO has been reported from leaves of individual corn plants exposed to 0.1–0.3 ppb NO (Hereid and Monson, 2001), from leaves of California live oak exposed to air containing NO (Delaria et al., 2018), from several nitrate-nourished plant species (Wildt et al., 1997) as well as pesticide-treated soybean leaves (Klepper, 1979). Additionally, recent plant physiological studies have started to reveal the mechanism of plant NO production and its importance for regulating growth and development, immunity, and signaling (Astier et al.,

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2017; R  , 2015; Yu et al., 2014), as well as for responding to pollutants and stress (Bison et al., 2018; Farnese et al., 2017; Velikova et al., 2008). In light of these advances, more targeted observations of foliar NO exchange probably should be conducted while taking these biological factors into consideration.

5 Summary and Conclusions

Using a branch enclosure technique and with controlled addition of trace gases, we obtained data on NO, NO₂, and O₃ leaf-level gas exchange from field experiments on several native tree species in a northern hardwood forest in Michigan, USA. To our knowledge, this is the first time such experiments have been done on North American tree species in a field study. The results provided a new dataset of NO_x and O₃ leaf-scale fluxes and have allowed comparisons of the gas exchange characteristics of mature trees compared to seedlings of these species in the lab and to mature European tree species in the field. The data also provide information, including an upper bound on NO₂ compensation points for these trees, to models of NO_x and O₃ dynamics at the canopy level, particularly for the forest at the PROPHET research site.

~~A brief survey of the foliar O₃ loss found that uptake by the deciduous trees also closely followed stomatal conductance, while the O₃ foliar deposition velocity for white pine was much larger than expected from leaf stomatal uptake alone. Removal via gas-phase chemical reactions was calculated to be negligible based on estimates of known BVOC emission rates and speciation, implying other non-stomatal pathways - cuticular uptake, dissolution to wet leaf surfaces, and/or chemical reactions at the leaf surface - are responsible for the additional ozone deposition, with their relative importance to be determined. A brief survey of the foliar O₃ loss found that the uptake by the deciduous trees also closely followed stomatal conductance, while the O₃ foliar deposition velocity for white pine was much larger than expected from leaf stomatal uptake alone, a difference that cannot be explained by leaf surface moisture conditions nor solely by gas-phase chemical removal. It is likely that other ozone removal pathways, such as cuticular uptake and the plant response to increased ozone level, were simultaneously responsible for this elevated deposition velocity.~~

The trace gas exchange characteristics of NO, NO₂, and O₃ at the leaf level varied depending on tree type and environmental conditions. For NO, there was no measurable foliar uptake from any of the trees studied here. On the contrary, there appeared to be a small emission of NO from white pine when NO was added to the enclosure. Leaf-level NO₂ uptake of bigtooth aspen and red oak was mainly through leaf stomata, with the leaf-level deposition velocity of NO₂ closely following predicted values based on the stomatal conductance of water and molecular diffusivity. The stomatal conductance of aspen was ~5 times higher than that of red oak (and thus the foliar NO₂ deposition velocity for aspen was also much higher). Because stomatal conductance is subject to a variety of factors including those intrinsic to plants, further investigation is needed to determine whether this difference is generally associated with the plant species or is environmentally driven. For white pine and red maple, the foliar NO₂ deposition velocity correlated with stomatal conductance, but there were additional factors that prevented ~~the~~ deposition from increasing as much as expected with ~~the~~ increasing conductance, suggesting the existence of internal

mesophyll resistance to uptake. Furthermore, for white pine, there was foliar NO₂ deposition when stomatal conductance was zero, suggesting a non-stomatal NO₂ loss pathway such as cuticular uptake.

The possible existence of an NO₂ compensation point was inferred by examining the linear relationship between NO₂ flux and ambient NO₂ concentration when the stomata were open, and the stomatal conductance was at least 60% of the maximum measured value. The results showed that the compensation point was ≤ 60 ppt for all trees and was statistically indistinguishable from zero within the measurement sensitivity. This finding does not support the suggested 1 ppb compensation point needed to reconcile the observed and model-simulated NO_x mixing ratios by Seok et al. (2013). Neither does it support any significant foliar NO₂ emission from these tree species at low ambient NO₂ conditions.

It is noteworthy that, beyond the findings in Seok et al. (2013), inclusion of an NO₂ compensation point not only provided the best agreement in terms of NO_x concentrations, but also gave the best agreement between simulated and observed atmosphere-biosphere NO_x fluxes at UMBS in summer 2016 (J. Murphy, personal communications, 2018). Evaluations of these simulations with [the Multi-Layer Canopy CHemistry Exchange Model \(MLC-CHEM\)](#), which was used in Seok (2013) have not yet included a direct comparison with the leaf-scale NO_x and O₃ fluxes reported here. Such a comparison could address both the observed large differences in the magnitude of the stomatal conductance for specific trees and its diurnal cycle, focusing on the early morning onset of stomatal opening and uptake. This would [help to establish further confirm](#) whether there is a leaf-scale NO_x emission flux due to an NO₂ compensation point, or if a strongly reduced NO_x uptake might partially explain the observed dynamics in the above- and in-canopy NO_x concentrations. This analysis would also benefit from more detailed temporally and vertically resolved NO_x concentration gradient observations compared to the Seok et al. (2013) study, which we measured in conjunction with the leaf-level work described here. This comparison ~~would be is~~ an essential next step [in attempting to potentially](#) reconcile the findings of this study with previous studies of NO_x exchange at the UMBS forest, and will be presented in a follow-up publication.

Our findings confirmed that the main conduit of trace gas foliar uptake is leaf stomata. A thorough grasp of the trace gas uptake efficiency hinges on an understanding of the leaf stomatal properties, which depend on the genetic makeup and developmental stage of the plant, as well as the environmental conditions of sunlight, water vapor, ambient temperature, soil, and nutrients. Meanwhile, the ~~characteristics of~~ additional factors affecting the foliar trace gas exchange, [such as](#) mesophyll resistance, cuticular uptake, and stress responses, are also subject to plant intrinsic and external conditions and remain to be better understood.

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Data availability. The data from this work are [are archived at https://umich.box.com/v/PROPHETAMOS2016](https://umich.box.com/v/PROPHETAMOS2016)~~available upon~~
~~request.~~

615 Author contributions. WW, SR, and JH constructed the enclosure chambers, deployed the instruments, and carried out the field
experiments. WW performed the data analysis and prepared the manuscript. LG and DH provided extensive comments and
suggestions for the manuscript. LG and DH initiated this project based on ~~their~~ previous fieldwork at UMBS and model
analysis regarding leaf-level gas exchange.

620 Competing interests. The authors declare that they have no conflict of interest.

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Figures and Tables

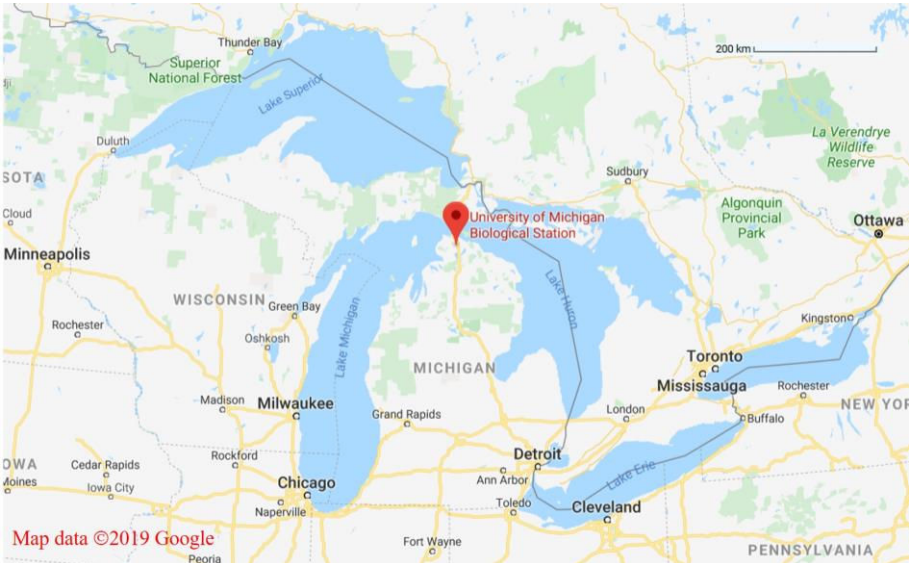
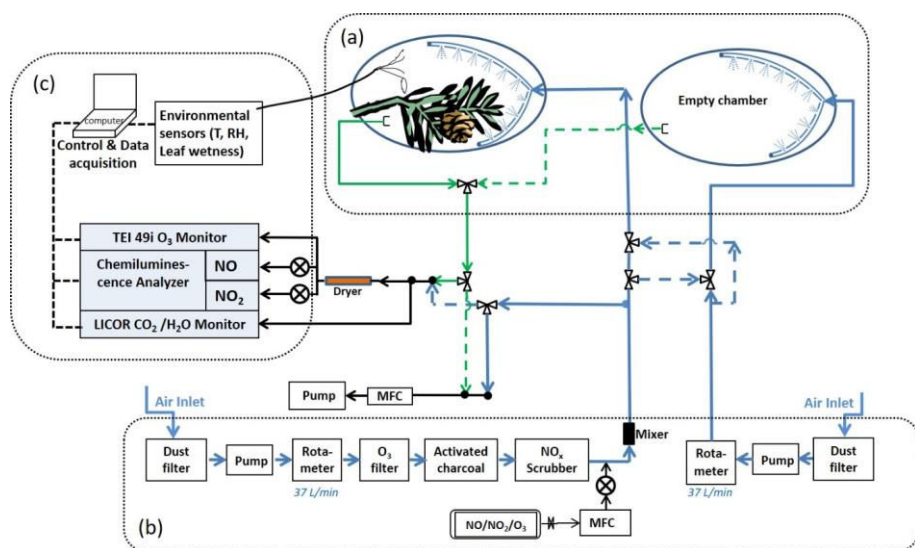


Figure 1. Location of the University of Michigan Biological Station, (45.56°N, 84.71°W), indicated by the red pin on the map. The map scale is shown in the upper right corner. (Map data ©2019 Google).

200 km



785 **Figure 2. Schematic of the enclosure experiment system. The system is comprised of three main parts as shown in the figure: (a) the**
enclosures, (b) the purge air flow system, and (c) the trace gas measurement instruments. The blue lines and arrows indicate the air
flowing into the enclosures; the green lines and arrows indicate the air flowing out of the enclosures; and the black lines and arrows
indicate the air sample flow and the balance flow (to maintain constant flow rates in the enclosures). NO and NO₂ gas standards
790 **were used for the controlled addition of these trace gases to the input air stream. Controlled O₃ addition was done by generating**
ozone on demand using a pen-ray lamp and ultra-high purity oxygen. See text for instrument details.

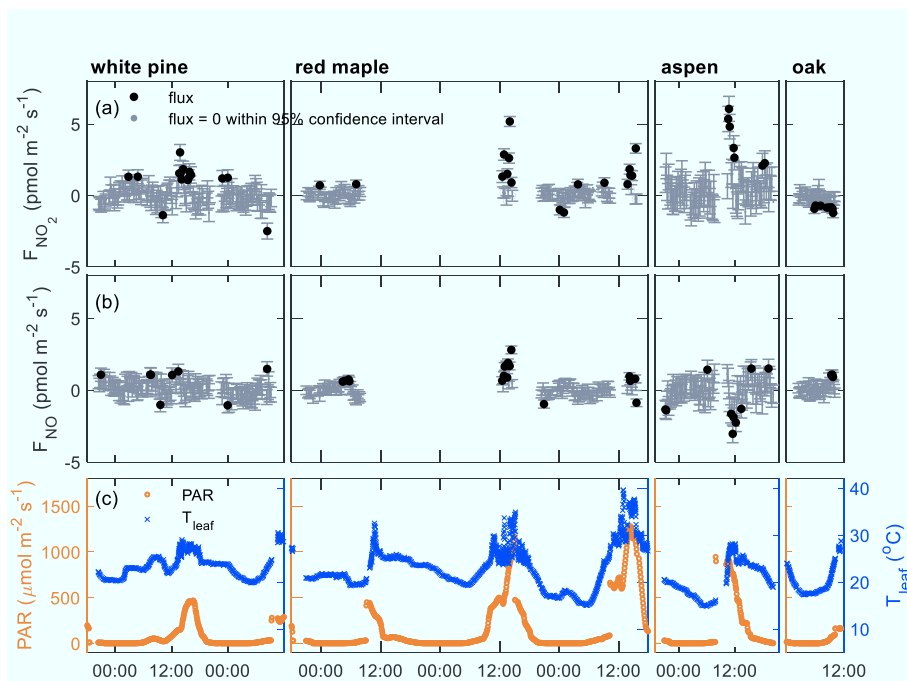


Figure 3. Apparent fluxes of NO₂ and NO when the enclosures were purged with scrubbed air. From left to right each panel corresponds to the enclosure of white pine, red maple, bigtooth aspen, and red oak. From top to bottom: (a) NO₂ flux, (b) NO flux, and (c) PAR (left axis, orange) and temperature of the enclosure leaves (right axis, blue). In panel (a) and (b), fluxes that are indistinguishable from zero within the 95% confidence interval are represented by gray dots; statistically significant fluxes are represented by black dots; error bars represent 1- σ measurements uncertainties propagated through the calculations.

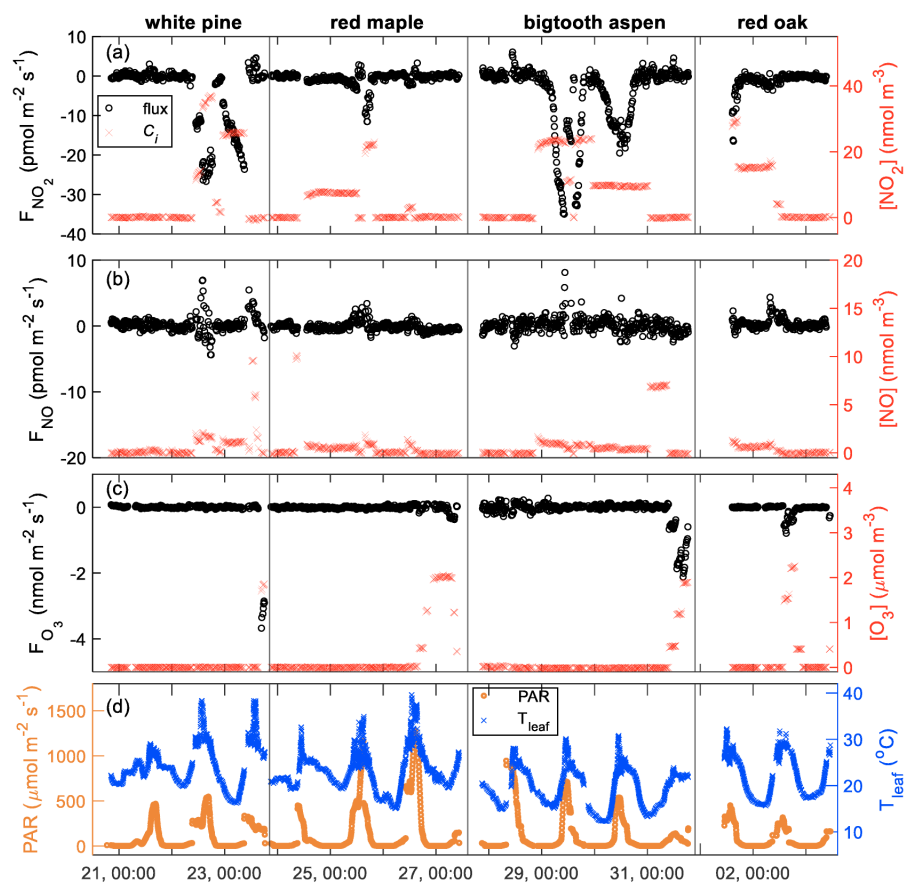


Figure 4. Time series of the enclosure gas exchange experiments from July 21 to August 3, showing the trace gas fluxes (black symbols: o, left axis) and input trace gas concentrations, C_i (red symbols: x, right axis) of, (a) NO_2 , (b) NO , and (c) O_3 . Solar irradiation PAR (orange: o, left axis) and temperature of the enclosure leaves (blue x, right axis) are shown in the bottom panel (d). The tree species are labeled for each enclosure period at the top of the figure. The x-axis tick label format is Day, HH:MM.

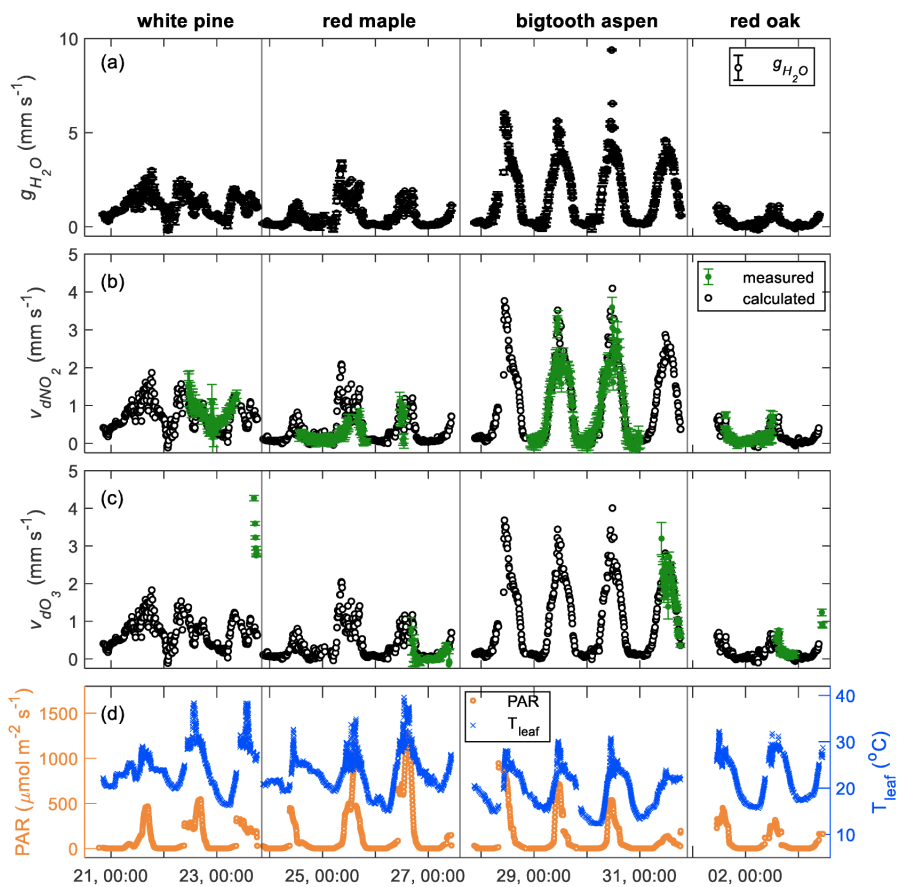


Figure 5. Time series plots of (a) the measured stomatal conductance of water; (b) the measured foliar deposition velocity (green symbols) and the calculated stomatal uptake rate (black symbols) of NO_2 ; (c) the measured foliar deposition velocity (green symbols) and the calculated stomatal uptake rate (black symbols) of O_3 ; and (d) the corresponding PAR and leaf temperature during the experiments from July 21 to August 3, 2016.

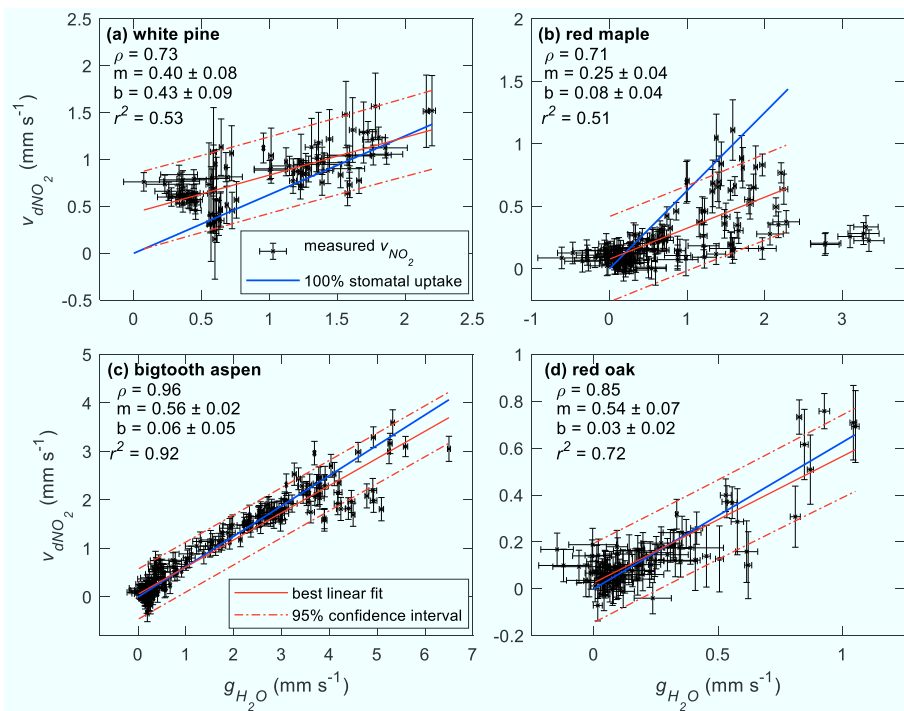


Figure 6: Scatter plots of foliar NO_2 deposition velocity ($v_{d\text{NO}_2}$) vs. stomatal conductance of water ($g_{\text{H}_2\text{O}}$) for (a) white pine, (b) red maple, (c) bigtooth aspen, and (d) red oak. The data points and their error bars are represented by the black symbols. The solid and dashed red lines are the best-fit linear regression and the 95% confidence bounds, respectively. The solid blue line shows the relationship between the deposition velocity and the stomatal conductance if NO_2 loss is entirely controlled by the stomata. The slope of the blue line is 0.62, the square root of the ratio of the molecular weight of water to NO_2 . Listed in each subplot under the tree name are the Pearson correlation coefficient (ρ), and the slope (m) and intercept (b) of the best fit linear regression line.

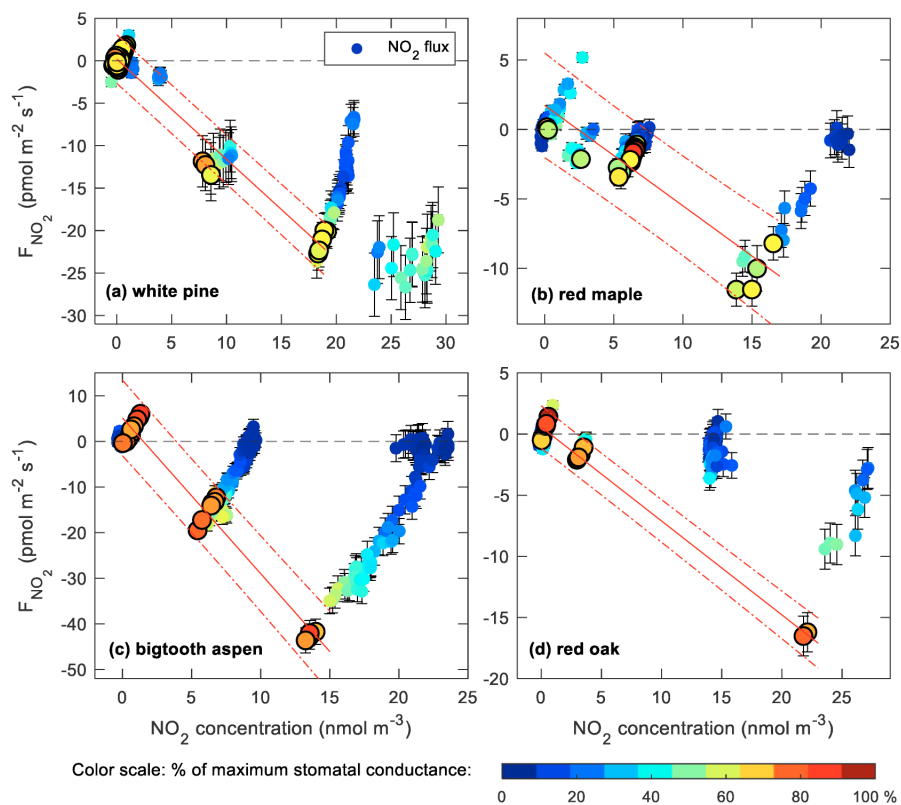


Figure 7. NO_2 flux toward the leaf surface vs. NO_2 concentration in each enclosure. Also shown is the stomatal conductance on a cool–warm color scale with the dark blue representing the lowest values and the red the highest values observed. Flux data with the stomatal conductance 60% of the observed maximum or higher were used for the linear extrapolation to find the compensation point. These data points are shown as larger symbols with a black outline/border. The solid and dashed red lines show the linear fit and 95% confidence bounds. The resulting compensation points are listed in Table 2.

830

Table 1. Detection limits of the foliar fluxes of NO₂, NO, and O₃. These were determined using the flux data obtained during the nighttime scrubbed air purges when no foliar gas exchange was expected. These limits mainly reflect the measurement precision of the trace gas concentrations at the inlets and outlets of the branch and reference enclosures, variations of the enclosure conditions over time, and the fluctuation of the purge air flow rate.

Detection limit	F_{NO_2} (pmol m ⁻² s ⁻¹)	F_{NO} (pmol m ⁻² s ⁻¹)	F_{O_3} (pmol m ⁻² s ⁻¹)
White pine	1.1	1.0	76.3
Red maple	0.6	0.6	68.0
Bigtooth aspen	2.0	1.3	233
Red oak	0.8	0.8	42.7

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Table 2. Compensation points from the flux vs. concentration linear fits in Fig. 7. The ranges of the stomatal conductance of the data used are also listed.

Tree species	White pine	Red maple	Bigtooth aspen	Red oak
Stomatal conductance (mm s ⁻¹)	1.8–3.0	1.8–3.5	4.5–6.0	0.8–1.3
Compensation point ± 95% confidence level (ppt)	4 ± 60	60 ± 119	38 ± 59	19 ± 56
Compensation point ± 95% confidence level (nmol m ⁻³)	0.2 ± 2.4	2.4 ± 4.9	1.6 ± 2.4	0.8 ± 2.3

Table 3: Comparison of foliar NO₂ deposition velocity from this work and earlier studies. The maximum velocity measured in each enclosure is listed with the corresponding light, RH, and leaf temperature at the time of the measurement.

Tree species	v_{NO_2} (mm s ⁻¹)	PAR (μmol m ⁻² s ⁻¹)	RH (%)	T _{leaf} (°C)	Source
Pinus strobus (white pine)	1.6	601	67	30	This work
Pinus strobus (white pine, seedling)	0.4	“Adequate to open leaf stomata”	n/a	29.4	Hansen (1989)
Acer rubrum (red maple)	1.1	1200	72	30	This work
Acer rubrum (red maple, seedling)	1.8	“Adequate to open leaf stomata”	n/a	29.4	Hansen (1989)
Quercus rubra (red oak)	0.76	1086	61	28	This work
Quercus agrifolia (California live oak)	1.23	1190	50–65	26	Delaria (2018)
Populus grandidentata (bigtooth aspen)	3.6	850	71	25	This work

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Author response to Anonymous Referee #1 of ACP-2020-149, “Measurement report: Leaf-scale gas exchange of atmospheric reactive trace species (NO₂, NO, O₃) at a northern hardwood forest in Michigan”

We greatly appreciate the thoughtful feedback provided by Anonymous Referee #1. The questions and comments have helped to improve and enhance the manuscript. Below, we address each comment individually. Referee comments are given in **Bold**, author responses are given in normal font, changes to make in the manuscript are given in blue.

P2 L60: The citation of Delaria et al., 2018 and reported estimation of 15-30% removal of soil-emitted NO_x is correct. Oak woodlands have a very low LAI. However, in Delaria and Cohen 2020 (now published and not in discussion), they report much larger canopy reductions for forests with more typical LAI, in line with the 25-55% loss previously reported.

The latest publication of Delaria and Cohen (2020) has been cited.

In the manuscript, P2:

“Implementing these results in a multi-layer single-column model, it was calculated that California oak woodland canopy removes 15-30% of soil-emitted NO_x, and other forests in California and Michigan, close to 60% (Delaria and Cohen, 2020).”

P3 L64: Extra parenthesis

It is removed.

P3 L83: Would be nice if the instrument were stated explicitly.

That information has been added to the text on P3. It was a chemiluminescent NO_x detector equipped with a highly NO₂ specific blue light converter.

P3 L96: Correct “folia” to “foliar”

Corrected.

P7 L205: Several studies have observed significant stomatal opening during the night (e.g. Dawson et al., 2007–10.1093/treephys/27.4.561). Consider adding a discussion of how, if this was occurring in your chamber, this assumption would have affected your results (if at all).

We have modified the manuscript text to address this point:

In the manuscript, P7:

“Nighttime transpiration in trees and shrubs has been measured in prior work, with reports of nighttime transpiration rates ranging from 0 to as much as 25% of the daytime value (Dawson et al., 2007), suggesting that leaf stomata may remain open at night for some plants. However, this possibility did not affect the above results as there was no evidence of a consistent concentration difference above zero between the enclosure outlet and inlet measurements.”

P8 L246: A more detailed description of your empty chamber photolysis corrections would be useful.

We have modified the paragraph to give a more specific and detailed description on the corrections based on the empty chamber measurements.

(Please also see response to RC#3, L245.)

In the manuscript, P8:

“There are a couple of factors that complicated the NO₂ gas exchange experiment. First, the NO₂ standard used for delivering NO₂ to the enclosure contains about 5% NO that was unavoidably added to the enclosure. Secondly, when there was intense direct sunlight, some NO₂ in the enclosure was photolyzed. While corrections for these interferences were done using the measurements from the reference enclosure, it is difficult to completely remove the artifact caused by NO₂ photolysis. This is because the sunlight exposure of the two enclosures, although situated side-by-side, was often uneven, and the measurements of the enclosures were done not simultaneously but sequentially. This problem is particularly pronounced for clear sky conditions with strong contrasts in sunlit and shaded conditions inside the canopy. The branch enclosure was always positioned to get more sun exposure than the reference enclosure if choices needed to be made. Therefore, the branch enclosure likely received more sunlight overall, even though it might be more shaded during some measurement cycles. Generally, for the periods of strong sunlight, there is residual NO after the correction against the reference enclosure is made. If we assume all this is due to an underestimation of NO₂ photolysis and make a further correction by combining the changes of NO and NO₂, the data quality is not improved while more noise is introduced to the data. Because of this and because we are not absolutely certain about all possible sources of NO_x from the branch enclosures, we prefer to adhere with the correction using only the reference enclosure measurements and view the resulting NO₂ flux as an upper bound, with possibly as much as 20% overestimation under direct sunlight conditions, which accounts for ~16% of all data during the NO₂ exchange experiments.”

P10 L314: Units for the intercept should be added. Additionally, under the resistance model framework you discuss, the relationship of V_d to g_{H2O} is non-linear. How might this affect your inferences of cuticular uptake?

We have added the units (mm s⁻¹) to the intercept.

For the second question:

In the resistance model framework at the leaf scale, the cuticular, stomatal, and mesophyll resistances are the main factors to determine the deposition velocity. Putting them together, the overall conductance (i.e. the inverse of the total resistance) to deposition is:

$$1/R_{\text{total}} = 1/R_{\text{cut}} + 1/(R_{\text{sto}} + R_{\text{meso}}).$$

Thus, if $R_{\text{cut}} \gg (R_{\text{sto}} + R_{\text{meso}})$ or R_{cut} is constant and $R_{\text{meso}} \ll R_{\text{sto}}$, the deposition velocity vs. stomatal conductance will be linear. If R_{meso} is significant compared to R_{sto} , the deposition velocity will be limited by R_{meso} as stomatal conductance increases. Of course, we don't know if and how R_{cut} and R_{meso} vary with environmental factors, which would potentially complicate the V_d to g_{H2O} relationship.

We added the following in the manuscript, P16 (of the marked-up copy):

“When extrapolated to zero stomatal conductance, the deposition velocity of NO₂ to white pine was 0.43 mm s⁻¹ (Figure 6a), implying deposition unrelated to leaf stomata, possibly to wet leaf surfaces and/or to leaf cuticula. This observation does not exclude the possible existence of these pathways when the stomata are open. A deposition velocity higher than expected based on the stomatal conductance would result if there is significant non-stomatal deposition. On the other hand, mesophyll resistance renders a lower deposition velocity than the expected value. There is no mechanistic reason why the deposition velocity associated with either a non-stomatal pathway or mesophyll resistance should remain constant or vary linearly with stomatal conductance. The relationship of deposition velocity, $v_{d_NO_2}$, and stomatal conductance, g_{H_2O} , would remain essentially linear as long as stomatal deposition dominates or the non-stomatal deposition term is constant while mesophyll resistance is small. However, if mesophyll resistance is significant, it would limit the increase of $v_{d_NO_2}$ with stomatal conductance. “

P12 L363: How high of emission rates would this require? Is it outside the range reported for trees of the species considered?

Yes, it does require an unreasonably high emission rate to match the amount of ozone removed.

Prompted by this comment, we have modified this argument using literature values of BVOC emission rates and speciation to give a more realistic estimation of the amount of ozone loss due to chemical reactions. It is < 1%.

(Please also see RC#3, p 3-4, “Can you comment more on O3 uptake...”).

In the manuscript, P13 (of the marked-up copy):

“Estimation of the possible contribution from gas-phase reactions with BVOCs was made as follows. The upper bounds of typical emission rates at 30°C and PAR level at 1000 μmol m⁻² s⁻¹ for monoterpenes and other BVOCs (excluding isoprene) are 3 and 5 μg C g⁻¹ h⁻¹, respectively (Guenther et al., 1994). The speciation of major BVOCs emitted by white pine at UMBS is based on Kim et al. (2011), including α- and β-pinene, limonene, linalool, α-humulene, and β-caryophyllene. Using the rate constants of the BVOCs with ozone reactions (Burkholder et al., 2015), and the residence time of 1.5 min in the enclosure plus ~6 sec in the sample line before reaching the detector, the estimated ozone loss due to gas-phase chemical reactions was less than 1%. Even with optimal light and temperature conditions for BVOC emission, the estimated gas-phase chemical removal would only be on the order of a few percent.”

P13 L395-396: should this be VPD?

Yes. It is fixed now.

P13 L403: “In the future,”? “In future work,”?

Suggestion is taken.

P14 L437: This was also a conclusion of Delaria et al., 2018.

We added this information in the text.

In the manuscript, P17 (of the marked-up copy):

“Delaria et al (2018) reached the same conclusion from their study on *Quercus agrifolia*.”

O3 deposition: There are a number of recent references discussing ozone deposition that are not included. The paper would be stronger if it placed itself in the context of these and other recent papers on the subject (e.g. Silva and Heald GRL 2018 <https://doi.org/10.1002/2017JD027278>, Kavassalis and Murphy GRL 2017 <https://doi.org/10.1002/2016GL071791> and Clifton et al. in Reviews of Geophysics 2020 <https://doi.org/10.1029/2019RG000670>).

We took the suggestion and added the references and relevant context in the manuscript.

In the manuscript, P3 (of the marked-up copy):

“Similarly, vegetation and plant surfaces also affect ozone levels through dry deposition (Clifton et al., 2019, 2020; Kavassalis and Murphy, 2017; Silva and Heald, 2018). In forested areas, ozone dry deposition occurs through leaf stomata as well as non-stomatal pathways including cuticular uptake, and wet or dry leaf surface reactions, while some O₃ is also removed by gas-phase chemical reactions e.g. with biogenic volatile organic compounds (BVOCs) and NO. Though these processes have been identified, the exact partitioning between the dry deposition pathways (and in-canopy chemical destruction) has not been unequivocally determined, hindering the ability to correctly assess ground-level ozone.”

Author response to Anonymous Referee #2 of ACP-2020-149: “Measurement report: Leaf-scale gas exchange of atmospheric reactive trace species (NO₂, NO, O₃) at a northern hardwood forest in Michigan”

We greatly appreciate the thoughtful feedback provided by Anonymous Referee #2. The questions and comments have helped to improve and enhance the manuscript. Below, we address each comment individually. Referee comments are given in **Bold**, author responses are given in normal font, changes made to the text in the manuscript are given in blue. The revised manuscript includes all the changes listed below.

L45: Maybe rather state “the air layers above the forest” instead of “free troposphere” as the air will first encounter the roughness sublayer or at nighttime the stable nocturnal boundary layer, then the mixed layer (daytime) or residual layer (nighttime) before reaching the free troposphere.

We have changed the phrase and now the text (page 2) reads:

“The relative differences in the time scales of the turbulent mixing and the chemical and physical sink processes determine the amount of NO_x removed within the canopy, with the remaining NO_x being released into the boundary layer.”

L96: folia => foliar

It’s corrected.

L132: To be able to judge potential surface effects and light absorption behavior (cutoff wavelength) of the enclosure please provide information about the material the bags were made of.

The material was Tedlar, polyvinyl fluoride. This info is now in the manuscript.

L140: Which material were the lines made of?

It’s polytetrafluoroethylene (PTFE), also added to the manuscript.

L147 and Fig. 2: Was there any special reasoning for putting the activated charcoal in front of the Purafil NO_x scrubber? Different to NO₂, NO is not well captured by charcoal, and therefore normally the Purafil is put in front of the charcoal as it not only adsorbs NO₂, but also oxidizes NO to NO₂ which is finally captured by the Purafil and the charcoal. At low ambient NO it might make no difference, but for higher NO this setup normally works better.

There was no specific reason to put the Charcoal before Purafil. Fortunately, during our work, the ambient NO concentrations were generally low in the relatively remote forest. Thank you very much for the information!

L163: Leaf wetness measurements are mentioned here but not presented or discussed in the manuscript. Could you please add these results? Alternatively, at least mention why they were not used.

Thank you for this point. We only used the leaf wetness results for qualitative/diagnostic purposes. Therefore, it was not mentioned in the data analysis. We added a short explanation in the text.

In the manuscript, P6:

“leaf wetness (for qualitative assessment of leaf conditions only) (S-LWA, Onset)”

L168: 1 min zero air measurement plus 2.5 min NO and 2.5 min NO₂ results in a 6 min cycle? Please clarify.

It was a mistake. It should be “1 min zero air measurement plus 2 min NO and 2 min NO₂”. Thank you for catching it.

L171: To have higher NO₂ absorption and less HONO photolysis the light emissions of the diode should be > 390 nm. The 365 nm is close to the HONO absorption band at 368 nm (Stutz et al., 2000). Nevertheless, even at the peak absorption of HONO the absorption cross section of HONO is about a factor of 1.5 smaller and under environmental conditions the HONO to NO₂ ratio is normally below 10 % (some cases up to 30 %). Therefore, the HONO interference in ambient air should be small. Due to the high surface to Volume ratio of the chambers, the HONO to NO₂ ratio could be higher and might depend on the chamber material. So please also provide information on the chamber material (see above comment).

We apologize for this mistake. We used Hamamatsu L11921-500 LED light sources. The peak wavelength was 385+/-5 nm.

L300: “Unknown measurement issue for water concentration”. As stated in the paper in line 315 the chamber air was not condensing according to the calculated dew points. Could the reason be instead evaporation of surface water films that form at a RH > 50% due to the deliquescence of deposited salts and other processes (e.g. Burkhardt and Eiden, 1994), Burkhardt and Hunsche, 2013).

Please see the response after the comment for L353.

L317: Liquid surface films can be formed by processes other than condensation (see esp. Burkhardt and Hunsche, 2013). Furthermore, a recent laboratory chamber study investigated the influence of liquid films (at RH below condensation) on deposition of Peroxyacetyl nitrate and O₃ to plants ((Sun et al., 2016) in detail.

Please see the response below.

L353: As mentioned in the above comments there are processes other than pure condensation that lead to the formation of liquid films. What about the surface wetness measurements? Do they show any changes at RH between 50 and 100 % that could be associated to enhanced O₃ uptake?

Regarding the last three comments, we thank the reviewer for bringing this to our attention. We do not have observational results to address the point about microscopic water film but it certainly is an aspect to investigate in future experiments. We have modified the text to reflect the possibility of microscopic

water film on the leaf surface. The leaf wetness sensor likely did not have similar deposits to aid the formation of water film because the sensor was cleaned before being placed in the enclosure. (The leaves were not). The wetness sensor showed relatively dry conditions, i. e. lower readings during the O₃ experiment compared to those at early morning hours, and the RH was between 70-80%.

We modified the manuscript text to reflect this possibility. On page 16 (of the marked-up copy):

“To assess the role of wet leaf surfaces to non-stomatal deposition, we calculated the white pine enclosure dew point using the temperature and relative humidity data and compared it to the measured leaf temperature. The leaf temperature was always higher than the dew point during the experiments, excluding the possibility of a wet leaf surface from the condensation of pure water. However, a microscopic water film may nevertheless form at a relative humidity as low as 50% if there are hygroscopic deposits on the leaf surface (Burkhardt and Eiden, 1994; Burkhardt and Hunsche, 2013; Sun, 2016). The microscopic water film could potentially modify gas exchange rates of water-soluble trace gases in the air. Data from this work do not contain information that can be used to delineate the possibilities of trace gas dissolution into microscopic water films or cuticular uptake. Further investigations with appropriately designed experiments, better measurement precisions, longer observation time, and under different environmental conditions are necessary to delineate the various possible deposition pathways and their dependencies.”

L361: As the reactions will not stop at the chamber outlet please provide the total residence time from within the chamber (1.5 min) to the analyzer as well?

The residence time in the sample line was about 6 seconds. (1/4" OD, 1/8" ID, 100 feet). This info is now added to the manuscript.

L365: A very good description of the relevant loss processes for O₃ is provided by a recent review (Clifton et al., 2020).

Yes. We have now cited this reference in the revised text.

L464: Please revisit this statement in the light of the above comments that liquid films can form at RH > 50 % and therefore contribute to enhanced O₃ surface deposition (esp. Sun et al., 2016).

Yes, this statement was modified to include the possibility of a thin water film at low RH and above the dew point.

In the manuscript, P18 (of the marked-up copy):

“A brief survey of the foliar O₃ loss found that uptake by the deciduous trees also closely followed stomatal conductance, while the O₃ foliar deposition velocity for white pine was much larger than expected from leaf stomatal uptake alone. Removal via gas-phase chemical reactions was calculated to be negligible based on estimates of known BVOC emission rates and speciation, implying other non-stomatal pathways - cuticular uptake, dissolution to wet leaf surfaces, and/or chemical reactions at the leaf surface – are responsible for the additional ozone deposition, with their relative importance to be determined.”

Author response to Anonymous Referee #3 of ACP-2020-149: “Measurement report: Leaf-scale gas exchange of atmospheric reactive trace species (NO₂, NO, O₃) at a northern hardwood forest in Michigan”

We greatly appreciate the thoughtful feedback provided by Anonymous Referee #3. The questions and comments have helped to improve and enhance the manuscript. Below, we address each comment individually. Referee comments are given in **Bold**, author responses are given in normal font, changes made to the text in the manuscript are given in blue. The revised manuscript includes all the changes listed below.

Line 96, misspelled foliar

It's corrected.

Line 170, What is the conversion efficiency for the photolysis? If it was different from 100% that would show up as a difference in calibration factor for NO and NO₂. Those calibrations ought to depend on ambient O₃ and light. Can you include some comment on how calibrations depend on ambient condition. I don't doubt you have done it all correctly, but this information might give some additional insight on interpreting the data.

The reactive trace gases in ambient air were scrubbed prior to the purge air entering the enclosure. So, measurements of the NO_x added to the enclosure were not subject to O₃ interference.

For the calibrations, UHP zero air, an NO standard, and UHP O₂ (for making O₃) were used. These steps did not involve ambient air.

During the calibrations, NO was calibrated directly using an NO standard diluted with the UHP zero air. For the NO₂ calibration, the same diluted NO standard was titrated with some O₃ to form a mixture of NO₂ and NO, then the NO was measured, which gave **the expected NO₂** counts (NO₂ = NO_{original_amount} – NO_{after_adding_O3}). Afterward, NO₂ was photolyzed to NO using the LED, and the total NO was measured, yielding **the measured NO₂** counts. The ratio: (**measured_NO₂** / **expected_NO₂**), is the conversion efficiency. It was fairly steady throughout the campaign at ~0.68. The instrument was calibrated every 7 hours.

We added in the manuscript, P6:

“The NO₂ to NO conversion efficiency was ~0.68. Because the ambient air was scrubbed to remove O₃ (and other trace gases) before entering the enclosures, the effect of ambient O₃ on NO_x measurements was negligible.”

Line 245; The discussion about quantifying the impact of NO₂ photolysis could be clarified a little more. Is your point that because the amount of NO emitted is poorly constrained you cannot just compute the uptake of NO_x by adding together NO and NO₂? Are there enough calibrations at different times of day to examine how the apparent NO₂ conversion efficiency varies with light level and account for conversion? What about conversion of NO to NO₂ by ambient O₃? This ought to be apparent by evaluating the variation in NO calibration constant as a function of O₃.

We have modified the text to explain the background correction using the empty enclosure better. Regarding photolysis, the two enclosures could not be measured simultaneously with one instrument, and they were in mid-canopy instead of under direct sunlight without branches above. So, the light exposure of the two chambers was not identical, however close, making the correction of NO₂ photolysis an approximation.

Conversion of NO₂ to NO was done in the instrument using an LED light source. The calibration was done every 7 hours.

The ambient air was scrubbed free from O₃ before it was sent to the enclosures. If there's any residual O₃, it's below the detection limit of the O₃ detector (<0.5 ppb). This is not enough O₃ to significantly convert NO to NO₂ within the residence time in the enclosure and the sample line (total ~96 sec).

(Please also see response to RC#1, P8 L246.)

The text in the manuscript is modified as follows, on page 8:

“There are a couple of factors that complicated the NO₂ gas exchange experiment. First, the NO₂ standard used for delivering NO₂ to the enclosure contains about 5% NO that was unavoidably added to the enclosure. Secondly, when there was intense direct sunlight, some NO₂ in the enclosure was photolyzed. While corrections for these interferences were done using the measurements from the reference enclosure, it is difficult to perfectly remove the artifact caused by NO₂ photolysis. This is because the sunlight exposure of the two enclosures, although situated side-by-side, was often uneven, and the measurements of the enclosures were done not simultaneously but sequentially. This problem is particularly pronounced when it was cloudless with a strong contrast of light and shade inside the canopy. The branch enclosure was always placed for better sun exposure than the reference enclosure if choices needed to be made. Therefore, the branch enclosure likely received more sunlight overall, even though it might be more shaded during some measurement cycles. Generally, for the periods of strong sunlight, there is residual NO after the correction against the reference enclosure is made. If we assume all this is due to an underestimation of NO₂ photolysis and make a further correction by combining the changes of NO and NO₂, the data quality is not improved while more noise is introduced to the data. Because of this and because we are not absolutely certain about all possible sources of NO_x from the branch enclosures, we prefer to adhere with the correction using only the reference enclosure measurements and view the resulting NO₂ flux as an upper bound, with possibly as much as 20% overestimation under direct sunlight conditions, which accounts for ~16% of all data during the NO₂ exchange experiments. “

Line 310, Can you go the next step after concluding that there is some mesophyll resistance? The effect of having a non-zero mesophyll resistance is a non-zero concentration inside the leaf. Using equation 2 and 3 you could solve for NO₂ concentrations internal to leaf. Similarly, for situations with excess deposition you could compute a value for cuticular deposition from the residual after subtracting the stomatal uptake.

Granted you wouldn't get a unique solution if there were both cuticular uptake and non-stomatal deposition. But you can make this section stronger by quantifying some values for the other processes

you point to. At the end of paragraph, having some values for range of mesophyll conductance would be better than just stating further investigation is needed.

The concentration of NO₂ in the leaf internal air space indeed can be calculated using equations 2 and 3. However, I am not sure if the numbers can be put into a meaningful context because the mesophyll resistance was inferred from data where the deposition rate was lower than the theoretical value from g_{H_2O} , but not systematically investigated. Also, if both cuticular uptake and mesophyll resistance exist, their effects would be in the opposite direction. NO₂ goes through disproportionation when it dissolves in water, thus the NO₂ pressure in the internal air space needs to be in steady state with the solution. These make it difficult to interpret the apparent internal NO₂ pressure.

Line 318 In addition to there being a possibility cuticular adsorption accounts for extra NO2 deposition you might also note that stomatal enclosure might not be complete. Discussion about whether stomatal conductance goes to zero shows up mostly in discussions seeking to explain sap flow or water flux that doesn't go to zero at night. It might not be as much of an issue for daytime periods, but could be noted just for completeness.

We did observe some water flux at nighttime, especially in the white pine enclosure. Correspondingly there was nighttime deposition of NO₂ to pine needles. The additional deposition attributed to cuticular or other non-stomatal processes was inferred by extrapolating the deposition velocity to zero stomatal conductance. The nighttime stomatal conductance of other trees was very small or close to zero.

Can you also comment on how much of the data are for conditions that the vapor pressure differences between leaf and ambient air are quite small so that stomatal conductance computation has larger uncertainty. At the limit when ambient air approached saturation and leaf and air temperature were equal the stomatal conductance couldn't be determined from water flux. You have noted in that leaf temperature always exceeded dewpoint in the context of discounting possibility of dew, but it is also relevant for evaluating how well stomatal conductance is defined.

About 10% of the data show small VPD (< 1 standard deviation from zero) under the conditions you mentioned. It usually happened in the early morning just before and around sunrise. The trace gas fluxes were relatively small during these times.

The above is added to the manuscript, P10 (of the marked-up copy):

“When the conditions are such that the difference between the leaf and air temperatures is small and the enclosure humidity is high, the difference between $C_{H_2O_leaf}$ and $C_{H_2O_enclosure}$ is also reduced, increasing the uncertainty in g_{H_2O} . In our measurements, this happened mostly from dawn to sunrise, accounting ~10% of the total data points, where the $(C_{H_2O_leaf} - C_{H_2O_enclosure})$ was within one standard deviation from zero.”

Can you comment more on O3 uptake. I agree that reaction with VOC could be an important loss process for O3 in addition to reaction with foliage. Your point would be stronger, however, by providing the rate constants for the VOC typically associated with white pine not just pointing out one with the highest reaction rate as well as noting that the oak and aspen are known isoprene emitters.

The part regarding gas-phase reactions has been modified to use emission rates and speciation from literature to give a more realistic estimation.

(Please also see RC#1 P12 L363.)

In the manuscript, page 13 (of the marked-up copy):

“Estimation of the possible contribution from gas-phase reactions with BVOCs was made as follows. The upper bounds of typical emission rates at 30°C and PAR level at 1000 $\mu\text{mol m}^{-2} \text{s}^{-1}$ for monoterpenes and other BVOCs (excluding isoprene) are 3 and 5 $\mu\text{g C g}^{-1} \text{h}^{-1}$, respectively (Guenther et al., 1994). The speciation of major BVOCs emitted by white pine at UMBS is based on Kim et al. (2011), including α - and β -pinene, limonene, linalool, α -humulene, and β -caryophyllene. Using the rate constants of the BVOCs with ozone reactions (Burkholder et al., 2015), and the residence time of 1.5 min in the enclosure plus ~6 sec in the sample line before reaching the detector, the estimated ozone loss due to gas-phase chemical reactions was less than 1%. Even with optimal light and temperature conditions for BVOC emission, the estimated gas-phase chemical removal would only be on the order of a few percent.”

Line 392 Say something more about the difference in chamber temperature for the aspen branch compared to the other species. Is this because the outside air temperature was cooler also, or on account differences in radiation? Large differences in conditions between the species make comparisons among them more difficult. Can you say anything about whether the chamber conditions relative to outside conditions were different for the species. Finally, the point that aspen have stomata on both sides of leaf ought to come first, and rather than speculate that they have more stomata per unit area find some data in the literature about this. There is no need for several lines of explanation about why the conductance is higher before this point about double sided stomata. The explanation only needs to explain the additional enhancement beyond twice. Likewise the other differences should come first. Does the extra water flux account for reduced temperatures in aspen enclosures?

The text has been modified based on the suggestions. The daily *average* temperatures inside and outside the chamber were similar. The air temperature was cooler on the days when aspen was sampled. The VPD during this time was smaller than in other enclosures when it was warmer.

We could not find in the literature specific records on the stoma distribution of aspen except populus in general. The part regarding stomata on both sides of aspen leaves was moved to the beginning of this part of the discussion as suggested.

In the manuscript, P14 (of the marked-up copy):

“Biological features may have contributed to this difference. Compared with the other three trees in this work, the aspen was younger and smaller. The enclosed branch was in the upper part of the crown containing developing new leaves. Past measurements, albeit on different species, have shown that for the same species under similar environmental conditions, leaves of young trees generally have higher stomatal conductance than old ones (Fredericksen et al., 1995; Hubbard et al., 1999; Niinemets, 2002; Yoder et al., 1994). Another possible reason for the observed high $g_{\text{H}_2\text{O}}$, while direct evidence has yet to be found, is the number of stomata. Many trees have stomata on only the lower (abaxial) leaf surface; however, trees that belong to the genus *Populus*, which includes aspen, are an exception. They have stomata on both sides (amphistomatous), a feature that allows increased photosynthetic rate and fast growth (Kirkham, 2014). If the bigtooth aspen leaves are indeed amphistomatous, a relatively high

g_{H_2O} can be expected. We compared the environmental conditions of the enclosures. The integrated PAR exposure levels were similar. The daily variation of the relative humidity in the bigtooth aspen enclosure was not significantly different from the others. In contrast, the average daily temperature was 19.2°C, cooler than the temperatures (23.9°C, 22.6°C, and 21.6°C) in the other enclosures, similar to the average ambient air temperature outside the enclosure during the same time, 19.1°C, and 23.6°C, 22.4°C, and 21.3°C. The combined conditions of moisture and temperature led to a relatively low vapor pressure deficit (VPD) in the aspen enclosure, 0.8 kPa, compared to 1.2 kPa (white pine), 1.0 kPa (red maple), and 1.4 kPa (red oak) in the others. Generally, VPD and g_{H_2O} are inversely correlated and a low VPD corresponds to a relatively high g_{H_2O} (Hubbard et al., 1999; Urban et al., 2017a, 2017b). However, because here we are comparing different tree species, we consider the observed results to stem from the combination of biological and environmental factors. Further examination of these factors is beyond the scope of this paper, nevertheless, it would be beneficial to take this temporal and spatial variability and inhomogeneity into account in model parameterizations of trace gas dynamics since plant stomata are the main conduit of NO_x and O₃ deposition over vegetation.”

Line 510: It would be preferable to have data availability point to an existing data set already available rather than just making it available on request. It easier for the investigator to just prepare the files once and submit to a data server (doesn't UMBS have this for work at the site). Additionally, if the investigators move or retire then data on request gets hard to find years from now.

Yes, the data will be archived at <https://umich.box.com/v/PROPHETAMOS2016>.