1 Measurements of carbonyl compounds around the Arabian Peninsula:

2 overview and model comparison

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11 Abstract

12 Volatile organic compounds (VOCs) were measured around the Arabian Peninsula using a research vessel during the AQABA 13 campaign (Air Quality and Climate Change in the Arabian Basin) from June to August 2017. In this study we examine carbonyl 14 compounds, measured by a proton transfer reaction mass spectrometer (PTR-ToF-MS), and present both a regional concentration 15 distribution and a budget assessment for these key atmospheric species. Among the aliphatic carbonyls, acetone had the highest 16 mixing ratios in most of the regions traversed, varying from 0.43 ppb over the Arabian Sea to 4.5 ppb over the Arabian Gulf, 17 followed by formaldehyde (measured by Hantzsch monitor, 0.82 ppb over the Arabian Sea and 3.8 ppb over the Arabian Gulf) and 18 acetaldehyde (0.13 ppb over the Arabian Sea and 1.7 ppb over the Arabian Gulf). Unsaturated carbonyls (C4-C9) varied from 10 19 to 700 ppt during the campaign, and followed similar regional mixing ratio dependence as aliphatic carbonyls, which were 20 identified as oxidation products of cycloalkanes over polluted areas. We compared the measurements of acetaldehyde, acetone and 21 methyl ethyl ketone to global chemistry-transport model (EMAC) results. A significant discrepancy was found for acetaldehyde, 22 with the model underestimating the measured acetaldehyde mixing ratio by up to an order of magnitude. Implementing a 23 photolytically driven marine source of acetaldehyde significantly improved the agreement between measurements and model, 24 particularly over the remote regions (e.g. Arabian Sea). However, the newly introduced acetaldehyde source was still insufficient 25 to describe the observations over the most polluted regions (Arabian Gulf and Suez), where model underestimation of primary 26 emissions and biomass burning events are possible reasons.

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36 1 Introduction

37 Carbonyl compounds (aldehydes and ketones) can be released into the air directly from a variety of primary biogenic and 38 anthropogenic sources. These include biomass burning (Holzinger et al., 1999;Holzinger et al., 2005;Koss et al., 2018), fossil fuel 39 combustion (Reda et al., 2014;Huang et al., 2018) including vehicles (Erickson et al., 2014;Dong et al., 2014), industrial solvent 40 use (Kim et al., 2008), and natural sources including plants and plankton (Zhou and Mopper, 1997;Warneke et al., 1999; Jacob et 41 al., 2002;Fall, 2003;Williams et. al., 2004; Bourtsoukidis et al., 2014). However, secondary production via the atmospheric 42 oxidation of hydrocarbons is considered to be more important for many carbonyl compounds including acetone and acetaldehyde 43 (Jacob et al., 2002; Millet et al., 2010).

44 Carbonyls have several important roles in the atmosphere. They form as stable intermediates directly after hydrocarbon oxidation 45 by hydroxyl radicals, O_3 or NO_3 , when the peroxy radicals initially formed react with each other (permutation reactions) or with 46 NO. Their production is linked to tropospheric ozone formation (Carlier et al., 1986) and their loss, through oxidation and 47 photolysis, is an important source of hydroxyl and hydroperoxyl radicals (HO_x) in the dry upper troposphere (Colomb et al., 2006). 48 Carbonyls serve as precursors of peroxyacetyl nitrates (PANs) which are important atmospheric NO_x (NO and NO₂) reservoir 49 species (Finlayson-Pitts and Pitts, 1997;Edwards et al., 2014;Williams et al., 2000). Carbonyl compounds are also important for 50 the growth of atmospheric particles (Kroll et al. 2005) thereby indirectly impacting the Earth's radiative balance. The atmospheric 51 lifetimes of carbonyl compounds varies considerably, from less than one day for acetaldehyde (Millet et al., 2010) to more than 15 52 days for acetone (Jacob et al., 2002; Khan et al., 2015) in terms of tropospheric mean lifetime. A multiday lifetime means that 53 carbonyl compounds can impact the air chemistry on local, regional and even hemispheric scales. The numerous primary and 54 secondary sources of carbonyl compounds, as well as their multiple loss routes (photolysis, OH, NO₃ and O₃ oxidation) makes 55 budget assessments difficult.

56 The most predominant atmospheric carbonyl compounds besides formaldehyde are acetaldehyde and acetone. They have been 57 reported to vary from tens or hundreds of ppt in remote areas (Warneke and de Gouw, 2001; Wisthaler, 2002; Lewis et al., 58 2005; White et al., 2008; Colomb et al., 2009; Read et al., 2012; Sjostedt et al., 2012; Tanimoto et al., 2014; Yang et al., 59 2014;Hornbrook et al., 2016; Wang et al., 2019) to several ppb in urban and polluted areas (Dolgorouky et al., 2012;Guo et al., 60 2013; Stoeckenius and McNally, 2014; Koss et al., 2015; Sahu et al., 2017; Sheng et al., 2018). Generally, secondary photochemical 61 formation from various precursors is the main source for those carbonyl compounds. However, several recent studies have shown 62 that acetaldehyde mixing ratios in both the remote marine boundary layer and the free troposphere could not be explained by 63 known photochemistry as implemented in various atmospheric chemistry models, which consistently underestimated the 64 measurements by an order of magnitude or more (Singh et al., 2003;Read et al., 2012;Wang et al., 2019). Several potential 65 additional acetaldehyde sources have been proposed including new hydrocarbon oxidation mechanisms, aerosol related sources 66 and oceanic sources. One possible source of acetaldehyde in the remote marine boundary layer is oceanic emission from the photo 67 degradation of colored dissolved organic matter (CDOM) in sea-surface water, where acetaldehyde could be produced together 68 with other low molecular weight carbonyl compounds (Kieber et al., 1990;Zhou and Mopper, 1997;Sinha et al., 2007;Dixon et al., 69 2013). Nevertheless, due to both limited airborne and seawater measurements of acetaldehyde, the importance of oceanic emission 70 is still under debate (Millet et al., 2010; Wang et al., 2019). In order to better understand the atmospheric budgets of acetaldehyde 71 (and the other carbonyl compounds), it is informative to analyze a dataset of multiple carbonyl compounds in both polluted and 72 clean environments, with influence from marine emissions, varying particulate loadings, and high rates of oxidation as shown in 73 Figure 1, which demonstrates the main formation pathways of acetaldehyde during this campaign.

- During the shipborne research campaign AQABA, carbonyl compounds were continuously measured by PTR-ToF-MS onboard a research vessel that circumnavigated the Arabian Peninsula. During the campaign, chemically distinct air masses were sampled, which had been influenced by primary emissions of hydrocarbons and inorganic pollutants (NO_x, SO₂ and CO) from petroleum industries and marine transport (Bourtsoukidis et al., 2019;Celik et al., 2019), by pollution from urban areas (Pfannerstill et al., 2019), and clean marine influenced air (Edtbauer et al., 2020). It is a unique dataset of carbonyl compounds encompassing starkly different environmental conditions from a region with few (or none) available in-situ measurements to date.
- 80 In this study, we provide an overview of carbonyl compound mixing ratios (aliphatic, unsaturated and aromatic) over the
- 81 Mediterranean Sea, Suez, Red Sea, Arabian Sea and Arabian Gulf. Using an empirical method based on measured hydrocarbon
- 82 precursors, we have analyzed the relative importance of the photochemical sources for the carbonyl compounds observed. The
- 83 analysis is then extended to include sources and transport by using a global model EMAC (5th generation European Centre –
- 84 Hamburg general model, ECHAM5 coupled to the modular earth submodel system, MESSy, applied to **a**tmospheric **c**hemistry).
- 85 Model measurement differences are investigated in both clean and polluted regions, with particular emphasis on acetaldehyde.



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- 87 Figure 1. Diagram of possible sources and formation pathways of acetaldehyde during the AQABA campaign.
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89 2 Methods

90 2.1 AQABA campaign

91 The AQABA campaign was conducted onboard the research vessel Kommandor Iona (KI) from the end of June to the end of 92 August 2017. The ship started from Southern France, proceeded across the Mediterranean, through the Suez Canal, around the 93 Arabian Peninsula into the Arabian Gulf and on to Kuwait, thereafter returning along the same route. Five laboratory containers 94 were loaded onto the vessel, containing multiple gas and particle phase measurement instruments as well as a weather station.

95 2.2 PTR-ToF-MS

96 2.2.1 Sampling and instrument set-up

A high-flow inlet (stainless steel tubing, 0.2 m diameter, 5.5 m tall and 3 m above the top of the containers and the front deck) was
installed at the front of the ship where the laboratory containers were located. A high flow of air (approximately 10 m³min⁻¹) was
drawn through the inlet which provided a common attachment point for sub-sampling lines for all gas-phase measurement
instruments. An air flow of 5 standard L min⁻¹ for the first leg and 3.5 standard L min⁻¹ for the second leg was pumped into the

on-board lab container through an ½" (O.D. = 1.27cm) FEP (fluorinated ethylene propylene) tubing (about 10 m long) insulated
and heated to 50-60 °C. A PTFE (polytetrafluoroethylene) filter was placed at the beginning of the inlet to prevent insects, dust
and particles entering the instruments. Every 2-5 days, the filter was replaced depending on the degree of pollution encountered.
Inside the VOC instrument container, the PTR-ToF-MS (8000, Ionicon Analytik GmbH Innsbruck, Austria) sampled a sub-flow
at 80-100 sccm through 1/8" (0.3175 cm) FEP tubing (~ 10 m in length, insulated and heated to 60 °C) from the main fast air flow
and then to the instrument's PEEK (polyether ether ketone) inlet which was likewise heated to 60 °C. The inlet system was shared
with total OH reactivity measurement (Pfannerstill et al., 2019).

108 The working principle of PTR-MS has been described in detail in previous studies (Lindinger et al., 1998; Ellis and Mayhew, 109 2013; Yuan et al., 2017). In brief, H_3O^+ primary ions are generated in the ion source, and then drawn into the drift tube where they 110 interact with sampled ambient air. Inside the drift tube, VOCs with a proton affinity greater than that of H_2O (691 kJ mol⁻¹) are 111 protonated by proton transfer from H_3O^+ . The resulting secondary ions are transferred to the detector, in this case a time-of-flight 112 mass spectrometer with mass resolution around 3500 for the first leg and 4500 for the second leg at mass 96amu. An internal 113 standard of trichlorobenzene ($C_6H_3Cl_3$) was continuously introduced into the instrument to ensure accurate mass calibration. Every 114 minute a spectrum with mass range (m/z) 0-450 was generated. The data reported in this study are all at 1 minute resolution unless 115 otherwise specified.

116 2.2.2 Instrument characterization

117 The instrument background was determined every three hours for 10 minutes with synthetic air. 4-point calibrations were 118 performed five times during the whole campaign using a standard gas mixture (Apel-Riemer Environmental inc., Broomfield, USA) 119 containing 14 compounds (methanol, acetonitrile, acetaldehyde, acetone, dimethyl sulfide, isoprene, methyl vinyl ketone, 120 methacrolein, methyl ethyl ketone, benzene, toluene, xylene, 1,3,5-trimethylbenzene and α -pinene). It has been previously reported 121 that the sensitivity of some compounds measured by PTR-MS are humidity dependent (de Gouw and Warneke, 2007). As the 122 relative humidity (RH) was expected to be high and varying (marine boundary layer with occasional desert air influence), humidity 123 calibration was combined with 4-point calibration by humidifying the gas mixture at different levels from 0% - 100% RH.

124 2.2.3 Data analysis

- The data were initially processed by the PTR Analyzer software (Müller et al., 2013) to identify and integrate the peaks. After
 obtaining the raw data (counts per second for each mass identified), a custom-developed python-based program was used to further
- 127 process the data to final mixing ratios. For compounds present in the standard gas cylinder, interpolated sensitivities based on the
- 128 five in-campaign calibrations were applied to derive the mixing ratios; while mixing ratios of the other masses were calculated by
- using a proton transfer reaction rate constant (k_{PTR}) of 2.0×10^{-9} cm³ s⁻¹. The uncertainty associated with the mixing ratios of the
- 130 calibrated compounds was around 6-17% (see Table S1). For the mixing ratios derived by assuming k_{PTR} , the accuracy was around
- 131 ±50% (Zhao and Zhang, 2004). The detection limit (LOD) was calculated from the background measurement with 3 times the
- standard deviation (3σ), 52 ± 26 ppt for acetaldehyde, 22 ± 9 ppt for acetone and 9 ± 6 ppt for methyl ethyl ketone (MEK) (Table
- 133 S1). Data below LOD were kept as determined for further statistic analysis (Figure 2 and Table 1).
- In this study, we have interpreted ion masses with the exact masses corresponding to $C_nH_{2n}O$, $C_nH_{2n-2}O$ and $C_nH_{2n-8}O$ as aliphatic, unsaturated and aromatic carbonyls, respectively (see exact protonated m/z in Table S2). Carbonyl compounds with a carbon number of three and above can be either aldehydes or ketones, which are not distinguishable with PTR-ToF-MS using H₃O⁺ as the
- 137 primary ion. However, laboratory experiments have shown that protonated aldehydic ions with carbon atoms more than three tend

to lose a H₂O molecule and fragment to other masses (Buhr et al., 2002;Spanel et al., 2002). Moreover, although both ketones and
aldehydes can be produced via atmospheric oxidation processes, ketones tend to have longer atmospheric lifetimes and higher
photochemical yields than aldehydes as mentioned in the introduction. The ratio of measured propanal to acetone was 0.07 in the
western Pacific costal region (Schlundt et al., 2017), 0.06 in an urban Los Angeles (Borbon et al., 2013) and 0.17 - 0.22 in oil &gas
production regions (summarized by Koss et al., 2017). Therefore, signals on the exact mass of carbonyl compounds from the PTRToF-MS are expected to be dominated by ketones, particularly in regions remote from the sources.

144 2.3 Meteorological data and other trace gases

145 The meteorological data were obtained by using a commercial weather-station (Sterela) which monitored wind speed, wind 146 direction, relative humidity (RH), temperature, speed of the vessel, and GPS etc. The actinic flux was measured by a spectral 147 radiometer (Metcon GmbH; Meusel et al., 2016). Non methane hydrocarbons (NMHC) mixing ratios were measured by a gas 148 chromatograph with flame ionization detector (GC-FID) online with the time resolution of 50 minutes. It measured hydrocarbons 149 (C2-C8) and aromatics (C6-C8) with the average LOD < 10 ppt for most of compounds. For a detailed instrumental description 150 see Bourtsoukidis et al. (2019). Formaldehyde mixing ratios were determined by a modified and optimized version of the 151 commercially available AL4021 (Aero-Laser, Germany), which utilizes the Hantzsch technique (Stickler et al., 2006). Methane 152 and carbon monoxide (CO) levels were monitored by a cavity ring-down spectroscopy analyzer (Picarro G2401). Ozone was 153 measured with an absorption photometer (Model 202 Ozone Monitor, 2B Technologies, Boulder, Colorado). Due to the potential 154 interference from sampling our own ship exhaust in which carbonyl compounds may be present (Reda et al., 2014), a filter was 155 applied to the data set based on the wind direction and NO_x , SO_2 and ethene levels.

156 2.4 Model simulations

157 The EMAC (ECHAM5/MESSy Atmospheric Chemistry) model was used to simulate atmospheric mixing ratios of several 158 carbonyl compounds along the cruise track covered during the AQABA campaign. The EMAC model is an atmospheric chemistry-159 general circulation model simulating the process of tropospheric air by considering processes which could influence trace gases 160 mixing ratios, such as transport, chemistry, interaction with ocean/land, dry deposition and so on (Pozzer et al., 2007;Pozzer et al., 161 2012;Lelieveld et al., 2016). The model applied in this study is a combination of the 5th generation of European Centre Hamburg 162 general circulation model (ECHAM5) (Roeckner et al., 2006) and the 2nd version of Modular Earth Submodel System (MESSy2) 163 (Jöckel et al., 2010), where a comprehensive chemistry mechanism MOM (Mainz Organic Mechanism) was deployed (Sander et 164 al., 2019). The model considers direct emissions (such as anthropogenic, biogenic, biomass burning etc.), atmospheric transport 165 and mixing, photochemical production of carbonyls (by OH, O_3 and NO_3), as well as physical and chemical removal processes. 166 The global fire assimilation system was used for biomass burning emissions (Kaiser et.al., 2012). The exchange of organic 167 compounds between ocean and atmosphere was considered in EMAC via the AIRSEA submodel, described in detail in Pozzer et 168 al. (2006). The transfer velocity is calculated online and the concentration in the water is prescribed by the user. For acetone, a 169 constant water concentration of 15 nmol/L is used, following the suggestion of Fischer et al. (2012). The model configuration in 170 the study is the same as the model applied in Bourtsoukidis et al. (2020), where a natural non-methane hydrocarbon source (ethane 171 and propane) was implemented. The model is in the resolution of T106L31 (i.e. $\sim 1.1^{\circ} \times 1.1^{\circ}$ horizontal resolution and 31 vertical 172 hybrid pressure levels up to 10 hPa) and the time resolution of 10 minutes. The measurement data of PTR-ToF-MS were averaged 173 to 10-minute resolution to match the model data resolution for further comparison.

175 3 **Results and discussion**

176 Around the Arabian Peninsula, the mixing ratios of individual carbonyl compounds varied over a wide range, from tens of ppt to 177 ppb levels. In this study, we divided the dataset geographically into eight regions (Figure 2, middle graph) to classify and 178 characterize the primary and secondary origins of carbonyl compounds. The regional delineations were: the Mediterranean Sea 179 (MS), Suez, Red Sea North (RSN), Red Sea South (RSS), Gulf of Aden (GA), Arabian Sea (AS), Gulf of Oman (GO) and Arabian 180 Gulf (AG), the same as those described by Bourtsoukidis et al. (2019). Figure 2 shows the abundance of aliphatic, aromatic and 181 unsaturated carbonyl compounds (carbonyls) for each region. Generally, aliphatic carbonyls were present at much higher mixing 182 ratios than aromatic and unsaturated carbonyls, with smaller carbonyl compounds (formaldehyde, acetaldehyde, C3 and C4 183 carbonyls) dominating the distribution. The mixing ratios of aliphatic carbonyls decreased dramatically from C5 carbonyls with 184 increasing carbon number. The box plots (Figure 2) also show that carbonyl compounds were measured at higher mixing ratios 185 and were more variable over Suez region and the Arabian Gulf. The abundance of carbonyl compounds varied markedly from 186 region to region with highest and lowest values found in the Arabian Gulf and the Arabian Sea, respectively. Table 1 shows the 187 mean, standard deviation and the median values for carbonyls in each region. In the following sections, each class of carbonyl 188 compounds are investigated in greater detail.



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Figure 2. Overview of mixing ratios for aliphatic, aromatic and unsaturated carbonyl compounds (CxHyO). The boxes represent 25% to 75% of the data with the central line and square indicating the median and the mean values, respectively. The whiskers show data from 5% to 95% and stars were drawn for the minimum and maximum data points within 1% to 99% of the dataset. 193 Within brackets under the region acronyms the main characteristics of the air masses are indicated, based on non-methane 194 hydrocarbon variability-lifetime results (b factor) from Bourtsoukidis et al. (2019). The data used for map plotting was from public

domain GIS data found on the Natural Earth web site (<u>http://www.naturalearthdata.com</u>) and was read into Igor using the IgorGIS
 XOP beta.

197 3.1 Aliphatic carbonyls (C_nH_{2n}O)

198 **3.1.1** Overview

199 Relatively high mean mixing ratios of aliphatic carbonyls were observed over the Arabian Gulf; the highest being acetone (C3 200 carbonyl compound) at 4.50 ± 2.40 ppb (median: 3.77 ppb), followed by formaldehyde at 3.83 ± 2.55 ppb (median: 3.02 ppb), 201 acetaldehyde at 1.73 ± 1.61 ppb (median: 1.02 ppb) and MEK (C4 carbonyl compound) at 0.87 ± 0.71 ppb (median: 0.56 ppb). As 202 the Arabian Gulf is highly impacted by the oil and gas industry, we compared the measurements of the four aforementioned 203 carbonyl compounds with those measured in the oil and gas region (Table 2). Expect for formaldehyde, acetaldehyde, acetone and 204 MEK were lower than the mixing ratios measured in the Uintah Basin, which was influenced by intensive oil and natural gas 205 activities (Koss et al., 2015). The general distribution of the aliphatic carbonyls in the Uintah Basin is similar to the Arabian Gulf, 206 with acetone levels being approximately twice as those of acetaldehyde. The carbonyl mixing ratios in the Arabian Gulf were 207 comparable to those measured in Hickory (PA, USA) surrounded by natural gas wells (Swarthout et al., 2015). Koss et al. (2017) 208 reported the max boundary layer enhancement of carbonyl compounds (C2-C7) measured during an aircraft measurement above 209 the most productive oil field in the United States (Permian Basin). Within the boundary layer of the Permian Basin, C5-C7 aliphatic 210 carbonyls had mixing ratios of 0.34 ppb, 0.08 ppb and 0.03 ppb; which are of the same magnitude but lower than the levels 211 measured over the Arabian Gulf for C5 (0.52 ± 0.48 ppb), C6 (0.19 ± 0.25 ppb) and C7 (0.04 ± 0.04 ppb) carbonyl compounds. 212 The sources of the major carbonyls in the Arabian Gulf will be discussed in details in section 3.1.2 and 3.4.3.

213 In contrast, aliphatic carbonyls had much lower average mixing ratios over the Arabian Sea and the Gulf of Aden especially for 214 C7-C9 carbonyls with mean mixing ratios below the detection limit for most of the time. During the summertime AQABA 215 campaign, the prevailing wind direction over the Arabian Sea was southwest (Figure S1). Four-day back trajectories indicate the 216 air was transported from the Arabian Sea (Northwestern Indian Ocean), passing East Africa coast, which brought relatively clean, 217 photochemically aged airmasses (Bourtsoukidis et al., 2019). The mean level of acetone over the Arabian Sea (0.43 ± 0.18 ppb, 218 median: 0.34 ppb) is close to the level measured in the marine boundary layer of Western Indian Ocean (0.49 ppb) (Warneke and 219 de Gouw, 2001) and comparable to other reported values from open-sea air measurement (see Table 2). Acetaldehyde was 220 measured at relatively low mixing ratios over the Arabian Sea (0.13 ± 0.12 ppb, median: 0.09 ppb), which is comparable than the 221 levels reported by the measurements done in northern-hemisphere open ocean (see, Table 2). Over the Gulf of Aden, acetaldehyde, 222 acetone and MEK had slightly higher mixing ratios than those over the Arabian Sea.

223 The Mediterranean Sea had somewhat higher levels of aliphatic carbonyls than the clean regions (the Arabian Sea and the Gulf of 224 Aden) but with acetone (above 2ppb) dominating the distribution. Much higher acetone level than acetaldehyde level was also 225 observed for some costal site measurement which was impacted by continental air (White et al., 2008; Schlundt et al., 2017, see 226 Table 2). Larger aliphatic carbonyls (C6-C9) were below the detection limit most of the time. The aliphatic carbonyls levels over 227 the Gulf of Oman were higher than the clean regions, while C1-C5 carbonyls were more variable over the Gulf of Oman compared 228 to those over the Mediterranean Sea. This is probably because the Gulf of Oman connects to the Arabian Gulf where intense oil 229 and gas industrial activities are located. Over the Gulf of Oman, polluted air from the nearby sources of the Arabian Gulf is 230 occasionally mixed with the clean air from the open sea (the Arabian Sea) under southeast wind conditions (Figure S1).

Another region where abundant aliphatic carbonyls were observed was Suez region. The air in this region was mainly influenced
 by nearby cities and marine transportation (ship emissions within the Suez Chanel) (Bourtsoukidis et al., 2019;Pfannerstill et al.,

2019). Therefore abundant precursors were available in Suez region, producing more carbonyls regionally especially for shorter-lived compounds (formaldehyde and acetaldehyde). Besides the local-scale emissions and photochemical production contribution to the carbonyls over Suez, the longer lived carbonyls (e.g. acetone) could be also transported from the Mediterranean Sea (where acetone was high). Four-day back trajectories indicate the air reaching Suez region was mostly originated from Europe continent passing over the Mediterranean Sea (Bourtsoukidis et al., 2019). Meanwhile, ocean uptake of acetone from the air due to polluted continental outflow (Marandino et al., 2005) as well as dilution and mixing with free tropospheric air during transport can modulate acetone mixing ratios. Although the mean mixing ratios of aliphatic carbonyls over Suez were much lower than those over the Arabian Gulf, the variations were still more significant than other regions (not including the Arabian Gulf, see Table 1).

Over the Red Sea, acetone was the most abundant aliphatic carbonyls followed by formaldehyde and acetaldehyde. The mixing ratios of acetaldehyde and acetone over the northern part of the Red Sea were similar to those levels measured in western Pacific coastal regions (South China Sea, Table 2). It is worth noticing that the levels of aliphatic carbonyls in the northern part of the Red Sea were almost two times higher than the southern part of the Red Sea. According to the four-day back trajectories reported by Bourtsoukidis et al. (2019), the measured air masses travelled to the northern part was from southern Europe and northeast Africa while the sourthern part was more influenced by air from the northern part of the Red Sea mixed with the air masses from desertic areas of central Africa. Therefore, less primary precursors as well as carbonyls were transported to the sourhtern part of the Red Sea compared to the northern part. Moreover, the unexpected sources of hydrocarbons (ethane and propane) from Northern Red Sea deep water reported by Bourtsoukidis et al. (2020) would lead to higher carbonyl levels in the Northen part compared with the Sourthern part due to the additional precursors in the Red Sea North. However, acetaldendye was still found to be significantly underestimated compared to the model results, even taking the deep-water source into consideration (section 3.3). This indicates that extra sources of acetaldehyde may exist, which will be disscussed in detail in section 3.4.

		Aliphatic Carbonyls								
		HCHO	CH3CHO	C3H6O	C4H8O	C5H10O	C6H12O	C7H14O	C8H16O	C9H18O
	mean	0.86	0.30	2.37	0.14	0.04	0.01	< LOD	< LOD	< LOD
MS	SD	0.41	0.25	0.37	0.05	0.02	0.01	n.a.	n.a.	n.a.
	median	0.80	0.25	2.32	0.12	0.03	0.01	< LOD	< LOD	< LOD
	mean	1.23	0.62	2.64	0.19	0.08	0.03	0.01	0.01	0.01
Suez	SD	0.76	0.58	0.77	0.15	0.08	0.02	0.01	0.01	0.01
	median	1.11	0.41	2.52	0.13	0.04	0.03	< LOD	< LOD	< LOD
	mean	0.99	0.50	2.17	0.27	0.12	0.04	< LOD	0.01	0.01
RSN	SD	0.78	0.26	0.45	0.11	0.07	0.02	n.a.	0.01	0.01
	median	0.73	0.46	2.17	0.25	0.10	0.04	< LOD	0.01	0.01
	mean	0.66	0.29	1.56	0.11	0.05	0.03	< LOD	0.02	0.03
RSS	SD	0.62	0.17	0.38	0.06	0.03	0.03	n.a.	0.03	0.05
	median	0.40	0.25	1.60	0.09	0.04	0.02	< LOD	< LOD	< LOD
	mean	0.69	0.15	0.81	0.04	0.02	0.03	0.01	< LOD	< LOD
GA	SD	0.33	0.09	0.27	0.02	0.01	0.02	0.01	n.a.	n.a.
	median	0.68	0.15	0.72	0.03	0.02	0.03	< LOD	< LOD	< LOD
	mean	0.82	0.13	0.43	0.02	0.01	0.01	< LOD	< LOD	< LOD
AS	SD	0.35	0.12	0.18	0.01	0.01	0.02	n.a.	n.a.	n.a.
	median	0.86	0.09	0.34	0.02	0.01	0.01	< LOD	< LOD	< LOD
	mean	1.27	0.25	1.23	0.10	0.08	0.04	0.01	0.01	0.01
GO	SD	0.59	0.12	0.40	0.06	0.04	0.03	0.00	0.01	0.01
	median	1.13	0.22	1.12	0.08	0.08	0.03	0.01	0.01	0.01
AG	mean	3.83	1.73	4.50	0.87	0.52	0.19	0.04	0.03	0.03
	SD	2.55	1.61	2.40	0.71	0.48	0.25	0.04	0.03	0.03
	median	3.02	1.02	3.77	0.56	0.31	0.10	0.03	0.02	0.02

280	Table 1. Mean, standard deviation (SD) and median mixing ratios of aliphatic, unsaturated and aromatic carbonyls in different
281	regions.

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		Aro	matic Carl	oonyls	Unsaturated Carbonyls						
		C7H6O	C8H8O	C9H10O	C4H6O	C5H8O	C6H10O	C7H12O	C8H14O	C9H16O	
MS	mean	0.02	0.01	< LOD	0.01	0.01	0.01	< LOD	< LOD	< LOD	
	SD	0.03	0.01	n.a.	0.02	0.01	0.01	n.a.	n.a.	n.a.	
	median	0.02	0.01	< LOD	0.01	0.01	0.01	< LOD	< LOD	< LOD	
	mean	0.09	0.03	< LOD	0.06	0.04	0.03	0.01	0.01	< LOD	
Suez	SD	0.20	0.04	n.a.	0.08	0.04	0.03	0.01	0.01	n.a.	
	median	0.02	0.01	< LOD	0.04	0.02	0.02	0.01	< LOD	< LOD	
	mean	0.09	0.05	0.02	0.03	0.03	0.04	0.02	0.01	0.01	
RSN	SD	0.10	0.06	0.02	0.02	0.03	0.04	0.02	0.01	0.01	
	median	0.06	0.04	0.01	0.02	0.03	0.03	0.01	0.01	< LOD	
	mean	0.05	0.04	0.03	0.01	0.02	0.02	0.01	0.03	0.01	
RSS	SD	0.06	0.06	0.03	0.01	0.02	0.02	0.01	0.07	0.01	
	median	0.01	0.01	0.03	0.01	0.01	0.01	< LOD	< LOD	< LOD	
	mean	0.02	0.02	0.01	0.01	0.01	0.02	0.01	< LOD	< LOD	
GA	SD	0.03	0.02	0.01	0.01	0.01	0.01	0.01	n.a.	n.a.	
	median	0.01	0.01	0.01	0.01	0.01	0.01	< LOD	< LOD	< LOD	
	mean	0.02	0.01	< LOD	0.01	0.01	0.01	< LOD	< LOD	< LOD	
AS	SD	0.03	0.01	n.a.	0.01	0.01	0.01	n.a.	n.a.	n.a.	
	median	0.01	0.01	< LOD	0.01	0.01	0.01	< LOD	< LOD	< LOD	
	mean	0.04	0.04	0.02	0.02	0.02	0.02	0.01	0.01	0.01	
GO	SD	0.06	0.05	0.03	0.01	0.01	0.01	0.01	0.01	n.a.	
	median	0.02	0.02	0.01	0.02	0.02	0.02	0.01	0.01	< LOD	
AG	mean	0.12	0.13	0.04	0.07	0.11	0.12	0.05	0.03	0.02	
	SD	0.14	0.10	0.04	0.06	0.10	0.10	0.05	0.03	0.02	
	median	0.08	0.10	0.03	0.04	0.07	0.09	0.04	0.03	0.01	

295 < LOD: the mixing ratios were lower than the limit of detection.

n.a.: not available

Locations	Lon./Lat.	Height	Time	Techniqu e	Formaldehyd e	Acetaldehyde	Acetone	MEK	Literature
Open sea		m							
Tropical Atlantic Ocean	10° N-0° N 35° W-5° E	18	Oct Nov.	PTR-MS	n.r.	n.r.	0.53	n.r.	(Williams et al., 2004)
Atlantic Ocean	50° N-50° S 10-60° W	18	Oct Nov.	PTR-MS	n.r.	0.18 (Northern H) 0.08 (Southern H)	0.6 (North) 0.2 (South)	n.r.	(Yang et al., 2014)
Western North Pacific Ocean	15-20∘ N 137∘ E	6.5 -14	May	PTR-MS	n.r.	n.r.	0.20-0.70	n.r.	(Tanimoto et al., 2014)
Western Indian Ocean	12° N-5° S 43-55° E	15	Feb Mar.	PTR-MS	n.r.	n.r.	0.49	n.r.	(Warneke and de Gouw, 2001)
Indian Ocean	19∘ N-13∘ S 67-75∘ E	10	Mar.	PTR-MS	n.r.	0.32-0.42 (continental outflow) 0.18-0.21 (equatorial marine)	1.11-2.08 (continenta l outflow) 0.51-0.62 (equatorial marine)	n.r.	(Wisthaler, 2002)
Southern Indian Ocean	30∘ S-49∘ S 30-100∘ E	15	Dec.	PTR-MS	n.r.	0.12-0.52	0.42-1.08	n.r.	(Colomb et al., 2009)
Costal Caribbean Sea	10-30∘ N 60-80∘ W	10	Oct.	HPLC	0.61	0.57	0.40	0.03	(Zhou and Mopper, 1993)
Cape Verde Atmospheric Observatory	16.86° N 24.87° W	10	2006 - 2011	GC-FID	n.r.	0.43 (0.19-0.67)	0.55 (0.23-0.91)	n.r	(Read et al., 2012)
Appledore Island, USA	42.97∘ N 70.62∘ W	5	Jul Aug.	PTR-MS	n.r.	0.40	1.5	0.20	(White et al., 2008)
Mace Head, Ireland	53.3∘ N 9.9∘ W	25	Jul Sep.	GC-FID	n.r.	0.44 (0.12-2.12)	0.50 (0.16-1.67)	n.r.	(Lewis et al., 2005)
Canadian Archipelago	68-75° N 60-100° W	Ship cruise	Aug. -Sep.	PTR-MS	n.r.	n.r	0.34	n.r	(Sjostedt et al., 2012)
Barrow Arctic	71.30° N 156.77° W	6	Mar. -Apr.	TOGA		0.10 ± 0.20	0.90 ± 0.30	0.19 ± 0.05	(Hornbrook et al., 2016)
South China Sea, Sulu Sea	2° N-15° N 108-124° E	10	Nov.	GC-MS	n.r.	0.86	2.1	0.06	(Schlundt et al., 2017)
Oil & Gas Horse Pool site, Uintah Basin, USA		Ground site	2012 - 2013	PTR-MS	3.71	4.27	7.97	2.81	(Koss et al., 2015)
Central United State		<600	Mar. - April	ToF- CIMS	1.13ª	0.5	1.5	0.2	(Koss et al., 2017)
Eagle Mountain Lake site, Texas, USA		Ground site	June	PTR-MS	n.r.	n.r.	3.2 (1.2-6.7)	0.3 (0.09-0.85)	(Rutter et al., 2015)

Hickory,	Ground site		PTR-MS		1.29	3.22	0.73	(Swarthout et
Donneylyoni		June		n r				
rennsyrvani				11.1.	(0.28 - 2.03)	$(1\ 45\ 4\ 99)$	(0.4 - 0.97)	al 2015)
a, USA					(0.28-2.03)	(1.45-4.77)	(0.4-0.97)	al., 2015)

307 n.r.: not reported in the literature.

308 a: formaldehyde was measured by laser-induced fluorescence (LIF)

309 3.1.2 Case studies of polluted regions: the Arabian Gulf and Suez

The primary emission sources in the Arabian Gulf and Suez regions are quite different. While the Arabian Gulf is dominated by oil and gas operations, Suez is more influenced by ship emissions and urban areas (Bourtsoukidis et al., 2019). Carbonyl compounds were most abundant in these two areas. For further insight, we focused on a time series of selected trace-gases and their inter-correlations to better identify the sources of the major aliphatic carbonyls. Meanwhile, we calculated the OH exposure ($[OH]\Delta t$) based on hydrocarbon ratios (Roberts et al., 1984;de Gouw et al., 2005;Yuan et al., 2012) for the polluted regions (Arabian Gulf and Suez) where primary emissions have been identified (Bourtsoukidis et al., 2019; Bourtsoukidis et al. 2020), to better understand the photochemical aging of the major carbonyls using the following equation:

317
$$[OH]\Delta t = \frac{1}{k_X - k_Y} \cdot (In \frac{[X]}{[Y]}\Big|_{t=0} - In \frac{[X]}{[Y]}),$$
 Eq. (1)

318 where X and Y refer to two hydrocarbon compounds with different rates of reaction with the OH radical (k). For this study, we 319 chose toluene (k_{OH+toluene}: 5.63E-12 cm³ molecule⁻¹s⁻¹) and benzene (k_{OH+benzene}: 1.22E-12 cm³ molecule⁻¹s⁻¹) (Atkinson and Arey, 320 2003), because both compounds were measured by PTR-ToF-MS at high frequency and these values showed a good agreement 321 with values measured by GC-FID (Figure S2). The approach detailed by Yuan et al. (2012) was applied to determine the initial emission ratio $\frac{[X]}{[Y]}\Big|_{t=0}$ in those two regions by only including nighttime data of benzene and toluene. We obtained initial emission 322 323 ratios (toluene to benzene ratio) of 1.38 for the Arabian Gulf and 2.12 for the Suez region. Koss et al. (2017) summarized the 324 toluene to benzene ratios observed in various locations and showed that urban and vehicle sources tend to have higher toluene to 325 benzene ratio (mean ~ 2.5) than the ratios of oil & gas sources (mean ~ 1.2). Therefore, the toluene to benzene ratios obtained for 326 those two regions agreed well with other studies done with similar emissions sources. The corresponding correlation plots of 327 toluene and benzene for those two regions can be found in Figure S3.



329

Figure 3. Case study of the Arabian Gulf. (a) Time series of selected species measured over the Arabian Gulf; (b) daytime
 correlation heat map of selected species; (c) nighttime correlation heat map of selected species.

332

333 Figure 3(a) shows the time series of acetaldehyde and acetone over the Arabian Gulf along with OH exposure ([OH]t) and ozone. 334 We further separated the data into daytime and nighttime and calculated correlations among the carbonyls and other selected 335 species (see Fig. 4b and c). Aliphatic carbonyls were well correlated with each other during the daytime and ozone had a generally 336 good correlation with C2-C7 carbonyls (r > 0.7) during the daytime but a much lower correlation during the night, indicating ozone 337 and carbonyls were co-produced via photochemical oxidation. Tadic et al. (2020) reported the net ozone production rate over the 338 Arabian Gulf (32 ppb d⁻¹) was the greatest over the Arabian Peninsula. They show that strong ozone forming photochemistry 339 occurred in this region, which would lead to abundant secondary photo-chemically produced products (including carbonyls). 340 However, it should be noted the good correlation between ozone and carbonyls could in part be due to carbonyls co-emitted with 341 ozone precursors (hydrocarbons) as primary emissions. In Figure 3 (a), the calculated OH exposure was high during the first night 342 in leg 1, where an elevation of acetone mixing ratio was observed while the mixing ratio of acetaldehyde remained relatively 343 constant. With limited OH radical abundance during the nighttime, the increased OH exposure indicates that the air reaching the 344 ship was photochemically processed (aged). Therefore, the increase of acetone was mainly from long-distance transport as acetone 345 has a much longer atmospheric lifetime than acetaldehyde. As the ship approached Kuwait, the calculated OH exposure was low 346 (starting from 7/30/2017, 12:00 am UTC), which is an indicator of nearby emission sources. The lifetime of the OH radical derived 347 from the measured OH reactivity also decreased from ~ 0.1 s to ~ 0.04 s during the same period (Pfannerstill et al., 2019). Oil 348 fields and associated refineries are densely distributed in the northwest of the Arabian Gulf region (United States Central 349 Intelligence Agency). The air reaching the ship when mixing ratios of acetone and acetaldehyde were highest was mainly from the 350 Northwest (Iraq oil field region) according to the back trajectories (Bourtsoukidis et al., 2019). This suggests that the air masses 351 encountered in Northwest Arabian Gulf were a combination of fresh emissions from nearby sources and photochemically processed

air transported from elsewhere. During the second leg, relatively low mixing ratios were identified in the same region (Northwest Arabian Gulf), which was mainly due to a greater influence of air masses originating from less populated desert regions of Northeast Iran (Bourtsoukidis et al., 2019) with much less influence from the oil field emissions, meaning less precursors were available for carbonyl production. Several plumes (extending over 2-3 hours) of elevated carbonyls with increased ozone were observed during the nighttime for both legs (Fig. 4a), indicating transport of highly polluted air.

357



358

Figure 4. Case study of Suez. (a) Time series of selected species measured over Suez; (b) correlation heat map of selected species
 during biomass burning plume (UTC 01:00 -06:00 August 24th 2017); (c) correlation heat map of selected species without the
 period of biomass burning plume.

362

363 For the Suez region (Gulf of Suez and Suez Canal), data were only available for the second leg. A significant increase of acetonitrile 364 (over 400 ppt) was observed just before entering the Great Bitter Lake (see Figure 4a), indicating an increasing influence of biomass 365 burning on the air composition (Lobert et al., 1990). Carbonyl compounds are important primary emissions in fresh biomass 366 burning plumes (Holzinger et al., 1999; Schauer et al., 2001; Holzinger et al., 2001; Koss et al., 2018) as well as being formed as 367 secondary products in more aged plumes (Holzinger et al., 2005). We further investigated the correlation coefficient among 368 carbonyls during the biomass burning plume (Figure 4b) in Suez. Carbonyls had a high correlation with acetonitrile, benzene and 369 among themselves, particularly for smaller carbonyls (acetaldehyde, C3-C5 carbonyls). The biomass burning emissions were 370 probably transported by on the prevailing northerly wind (Figure S1) above Northeast Egypt where crop residues especially rice 371 straw is often directly burned in the open fields (Abdelhady et al., 2014; Said et al., 2013; Youssef et al., 2009). Besides the direct 372 biomass burning emission, the high mixing ratios and the good correlations of carbonyls could also have resulted from other 373 sources as hydrocarbons (alkanes, alkenes and aromatics) which were elevated at the same time. Similar to conditions identified 374 over the Arabian Gulf, elevated OH exposure accompanied with increasing acetone mixing ratio was observed during the first 375 night over the Gulf of Suez, indicating aged air mass transportation. The OH exposure was then significantly lower during the 376 daytime, when mixing ratios of carbonyls and alkanes increased as well. This indicates the presence of emission sources nearby. Oil refineries located in the coastal side of Suez and oil tank terminals located in the northern part of the Gulf of Suez are likelysources.

sources.

379 3.2 Unsaturated and aromatic carbonyls (C_nH_{2n-2}O), (C_nH_{2n-8}O)

380 3.2.1 Overview

381 The mixing ratios of unsaturated carbonyls were generally ~ 10 ppt or lower than the LOD over the Mediterranean Sea and the 382 clean regions (the Arabian Sea and the Gulf of Aden). The Red Sea region and the Gulf of Oman had slightly higher levels (LOD 383 -40 ppt). The highest values were again observed in the Arabian Gulf (20 - 110 ppt) followed by Suez (LOD - 60 ppt). The 384 numbers represent the range of the mean mixing ratios of unsaturated carbonyls in each region. In terms of the mixing ratio 385 distribution (Figure 2), the peak value was usually observed at C5 or C6 unsaturated carbonyls over most regions except for Suez 386 where C4 carbonyl had the highest mixing ratio. Based on chemical formulas, unsaturated carbonyls can be either cyclic carbonyl 387 compounds or carbonyls containing a carbon-carbon double bond. Therefore, the air chemistry could differ considerably depending 388 on the compound assignment. A detailed analysis of the chemistry of the unsaturated carbonyls measured will be given in the 389 following section 3.2.2.

390 Regional variability was also observed for aromatic carbonyls with highest levels observed over the Arabian Gulf and Suez, and 391 much lower mixing ratios over the Arabian Sea, Mediterranean Sea and Gulf of Aden (Table 1). Several studies using PTR-MS 392 have reported values for m/z 107.049 (C7 aromatic carbonyls) attributed to benzaldehyde (Brilli et al., 2014; Koss et al., 2017; Koss 393 et al., 2018), m/z 121.065 (C8 aromatic carbonyls) attributed to tolualdehyde (Koss et al., 2018) or acetophenone (Brilli et al., 394 2014) and m/z 135.080 (C9 aromatic carbonyls) attributed to methyl acetophenone (Koss et al., 2018) or benzyl methyl ketone 395 (Brilli et al., 2014) or 3,5-dimethylbenzaldehyde (Müller et al., 2012). Atmospheric aromatic carbonyls are produced via 396 photochemical oxidation of aromatic hydrocarbons (Finlayson-Pitts and Pitts Jr, 1999; Wyche et al., 2009; Müller et al., 2012) and 397 benzaldehyde was reported as having primary sources from biomass burning and anthropogenic emissions (Cabrera-Perez et al., 398 2016). Around the Arabian Peninsula, the level of aromatic carbonyls declined with increasing carbon number over most of the 399 regions except in the Red Sea South, Gulf of Oman and Arabian Gulf where C7 carbonyls were comparable to C8 carbonyls (Figure 400 2). Interestingly, only in the Suez region, were the C7 aromatic carbonyls more abundant than other aromatic carbonyls, whereby 401 the mean value (90 \pm 200 ppt) was much higher than the median value (20 ppt), indicating strong primary sources of benzaldehyde 402 in Suez. Otherwise, toluene was found to be more abundant over Suez with mean mixing ratios of 271 ± 459 ppt than over other 403 regions (the mean over the Arabian Gulf: 130 ± 160 ppt) which would also lead to higher benzaldehyde as it is one of the OH-404 induced oxidation products of toluene via H-abstraction (Ji et al., 2017).

405 3.2.2 Potential precursors and sources of unsaturated carbonyls

406 Unsaturated carbonyls measured by PTR-MS have been only rarely reported in the atmosphere with the exception of methyl vinyl 407 ketone and methacrolein (C4 carbonyls) which are frequently reported as the oxidation products of isoprene (Williams et al., 2001; 408 Fan and Zhang, 2004; Wennberg et al., 2018). According to the GC-FID measurement, isoprene was below the detection limit for 409 most of the time during the AOABA cruise with the highest values observed in Suez (10 - 350 ppt). This shows that the AOABA 410 campaign was little influenced by either terrestrial or marine isoprene emissions. However, we observed unexpected high levels 411 on mass 69.070, which is usually interpreted as isoprene for PTR-MS measurements. Significant enhancements were even 412 identified while sampling our own ship exhaust (in PTR-MS but not GC-FID), suggesting the presence of an anthropogenic 413 interference at that mass under these extremely polluted conditions. Several studies have reported possible fragmentations of cyclic 414 alkanes giving mass (m/z) 69.070. These include: a laboratory study on gasoline hydrocarbon measurements by PTR-MS 415 (Gueneron et al., 2015), a GC-PTR-MS study of an oil spill site combined with analysis of crude oil samples (Yuan et al., 2014) 416 and an inter-comparison of PTR-MS and GC in an O&G industrial site (Warneke et al., 2014). From those studies, other 417 fragmentations from C5-C9 cycloalkanes including m/z 43, m/z 57, m/z 83, m/z 111 and m/z 125 were identified together with 418 m/z 69. Cyclic alkanes were directly measured in oil and gas fields (Simpson et al., 2010; Gilman et al., 2013; Li et al., 2017; Aklilu 419 et al., 2018), vehicle exhaust (Gentner et al., 2012; Erickson et al., 2014), vessel exhaust (Xiao et al., 2018), accounting for a non-420 negligible amount of the total VOCs mass depending on the fuel type. Koss et al. (2017) reported enhancement of cyclic alkane 421 fragment signals and increased levels of unsaturated carbonyls measured by PTR-ToF-MS over O&G region in the US. The 422 unsaturated carbonyls (C5-C9) were assigned as oxidation products of cycloalkanes. Therefore, we examined the correlations 423 between m/z 69.070 and other cycloalkane fragments over the Arabian Gulf and Suez, where anthropogenic primary emissions 424 were significant. As shown in Figure 5, m/z 83 was the most abundant fragment and it correlated better with m/z 69 than the other 425 two masses, strongly supporting the presence of C6 cycloalkanes (methylcyclopentane and cyclohexane). The other two masses 426 are distributed in two or three clusters, suggesting compositions of different cycloalkanes. M/z 43 and m/z 57 (fragments of C5 427 cycloalkanes) had lower correlations with other fragments (not shown in the graph) as they are also fragments of other higher 428 hydrocarbons. Thereby we could assign those unsaturated carbonyls as photochemical oxidation products (i.e. cyclic ketones or 429 aldehydes) from their precursor cycloalkanes.



Figure 5. Scatter plots of m/z 69.070 and other cycloalkane fragment masses over the (a) Arabian Gulf and (b) Suez region.

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433 As shown in Figure 2 and Table 1, C6 unsaturated carbonyls displayed higher mixing ratios than any other unsaturated carbonyls 434 over the Arabian Gulf while C5 unsaturated carbonyl was slightly higher than C6 in Suez. Bourtsoukidis et al. (2019) derived 435 enhancement ratio slopes from pentane isomers and established that the Arabian Gulf is dominated by oil and gas operations and 436 that Suez is more influenced by ship emissions. Therefore, as the Arabian Gulf had much more active O&G activities than Suez, 437 our findings agree with Koss et al. (2017) who showed that C6 unsaturated carbonyls should be more abundant than C5 carbonyls 438 since more precursors for C6 unsaturated carbonyls are emitted from active oil fields. It is worth mentioning that in Figure 5 (b) 439 one cluster at the bottom showed m/z 69.070 had no correlation with other three masses. Those points correspond to the time when 440 the GC measured significant elevated isoprene while passing through the narrow Suez Canal where some vegetation (e.g. palms 441 and some agriculture) was present close to shore, meaning m/z 69.070 during this period was isoprene. At the same time, m/z 442 71.049 (C4 unsaturated carbonyl) increased from 20 ppt to 220 ppt. Isoprene oxidation products (MVK and methacrolein) were 443 probably the major contribution to the C4 unsaturated carbonyls in this period. This also explains why C4 carbonyl dominated the 444 distribution of unsaturated carbonyls over Suez.

In the other regions (especially more remote areas), the cyclic alkane fragmentation masses had much lower abundance, leading to much less unsaturated carbonyls due to lack of precursors. Meanwhile, m/z 69.070 ($C_5H_8H^+$), m/z 83.086 ($C_6H_{10}H^+$) and m/z 97.101 ($C_7H_{12}H^+$) could also be fragmentations from corresponding aldehydes losing one water molecule as mentioned in section 2.3.3. Missing information of the chemical structure of unsaturated carbonyls and knowledge of their precursors, preclude detailed investigation of the sources of large unsaturated carbonyls in these areas.

450 **3.3** Model comparison of acetaldehyde, acetone and MEK

451 We compared our measurement results of acetaldehyde, acetone and MEK to those predicted by the global model "EMAC" 452 (ECHAM5/MESSy2 for Atmospheric Chemistry). From the results shown in Figure 6, the model predicted acetone much better 453 than acetaldehyde and MEK. In general, the model broadly captured the major features identified during the campaign such as 454 much higher levels of carbonyls mixing ratios over the Arabian Gulf and Suez and relatively low levels over the Arabian Sea. The 455 mean measurements-to-model ratios indicated that acetone was overestimated by a factor within 1.5 over the Arabian Sea, Gulf of 456 Aden and Gulf of Oman, and underestimated by a factor within 2.5 over the other regions. In contrast, the model underestimated 457 MEK within a factor of 4 over most of the regions except for the Gulf of Oman where MEK was overestimated (median values 458 were taken here as the mean values substantially deviated from the medians over Suez, Gulf of Oman and Arabian Gulf). The 459 model underestimation was most significant for acetaldehyde, which is underpredicted by a factor (median values) of more than 6 460 over the Red Sea North, ~ 4 over the Arabian Sea and Arabian Gulf and between 1 and 4 over other regions. A strong natural non-461 methane hydrocarbon source from deep water in the Northern Red Sea was implemented in the model (Bourtsoukidis et al., 2020). 462 Although the model representation of acetaldehyde and other carbonyls was clearly improved after including the deep water source 463 of ethane and propane (Figure S4), the underestimation of acetaldehyde was still significant over the Red Sea North as shown in 464 Figure 6(a), indicating further missing sources. For acetaldehyde and MEK, the discrepancy was also significant over the Arabian 465 Sea where acetone was in contrast, overestimated. Since acetaldehyde had the biggest bias from the model prediction, we further 466 investigate the possible missing sources of acetaldehyde.



468



473 3.4 Missing sources of acetaldehyde

In this section we investigate the following processes as potential sources of acetaldehyde: (1) production as an inlet artifact, (2)
oceanic emission of acetaldehyde, (3) anthropogenic primary sources, (4) biomass burning sources, and (5) other possible
secondary formation pathways.

477 3.4.1 Inlet artifact

478 Northway et al. (2004) and Apel et al. (2008) reported that heterogeneous reactions of unsaturated organic species with ozone on 479 the wall of the Teflon inlet can cause artifacts signal of acetaldehyde but not to acetone. During AQABA, the highest and the most 480 variable ozone mixing ratios were observed during the campaign over the Arabian Gulf (mean: 80 ± 34 ppb) and the Red Sea North 481 $(66 \pm 12 \text{ ppb})$, where a modest correlation was found between acetaldehyde and ozone over the Arabian Gulf (r²=0.54) and no 482 significant correlation over the Red Sea North ($r^2=0.40$). However larger correlation coefficients were identified between ozone 483 and other carbonyls over the Arabian Gulf (see Figure S5), which suggests that the correlation was due to atmospheric 484 photochemical production rather than artifacts. Moreover, acetaldehyde was found to have a much worse correlation with ozone 485 during the nighttime compared to the correlation during the daytime over the Arabian Gulf (Figure 3b and c), which also indicates 486 that inlet generation of acetaldehyde was insignificant. Over other regions, especially the remote area (the Arabian Sea and Gulf 487 of Aden), ozone was relatively constant and low, with poor correlation with acetaldehyde mixing ratios. Although we cannot 488 completely exclude the possible existence of artifacts, the interference is likely to be insignificant in this dataset.

489 3.4.2 Oceanic emission

490 A bias between measured acetaldehyde and global model simulations has been observed in previous studies conducted in the 491 remote troposphere (Singh et al., 2003; Singh, 2004; Wang et al., 2019) and in the marine boundary layer (Read et al., 2012). The 492 aforementioned studies emphasized the potential importance of the sea water acting as a source of acetaldehyde emission via air-493 sea exchange. No significant correlation was found between acetaldehyde and DMS, a marker of marine biogenic emission which 494 is produced by phytoplankton in seawater (Bates et al., 1992) (see Figure S6). This indicates that the direct biogenic acetaldehyde 495 emissions from the ocean are probably insufficient to explain the measured acetaldehyde. More likely, acetaldehyde and other 496 small carbonyl compounds can be formed in the sea especially in the surface microlayer (SML) via photodegradation of colored 497 dissolved organic matter (CDOM) (Kieber et al., 1990; Zhou and Mopper, 1997; Ciuraru et al., 2015). Zhou and Mopper (1997) 498 calculated the exchange direction of small carbonyls based on measurement results and identified that the net flux of acetaldehyde 499 was from sea to the air whereas formaldehyde was taken up by the sea. Sinha et al. (2007) characterized air-sea flux of several 500 VOCs in a mesocosm experiment and found that acetaldehyde emissions were in close correlation with light intensity (r=0.7). By 501 using a 3-D model, Millet et al. (2010) estimated the net oceanic emission of acetaldehyde to be as high as 57 Tg a^{-1} (in a global 502 total budget: 213 Tg a⁻¹), being the second largest global source. A similar approach was applied in a recent study done by Wang 503 et al. (2019), reporting the upper limit of the net ocean emission of acetaldehyde to be 34 Tg a⁻¹. Yang et al. (2014) quantified the 504 air-sea fluxes of several OVOCs over Atlantic Ocean by eddy covariance measurements, showing ocean is a net source for 505 acetaldehyde. Although Schlundt et al. (2017) reported uptake of acetaldehyde by the ocean from measurement-inferred fluxes in 506 western Pacific coastal regions, to our knowledge, there is no direct experimental evidence showing the ocean to be a sink for 507 acetaldehyde.

508 In order to test the importance of the oceanic emission of acetaldehyde, we implemented this source in EMAC model. The measured 509 sea water concentration of acetaldehyde was not available for the water area around the Arabian Peninsula. Wang et al. (2019) 510 estimated the global average acetaldehyde surface seawater concentrations of the ocean mixed layer using a satellite-based 511 approach similar to Millet et al. (2010), where the model estimation agreed well with limited reported measurements. From the 512 Wang et al. (2019) results, the averaged seawater concentration of acetaldehyde around Arabian Peninsula was generally much 513 higher from June to August. As the photodegradation of CDOM is highly dependent on sunlight, the air-sea submodel (Pozzer et 514 al., 2006) was augmented to include throughout the campaign a scaled acetaldehyde seawater concentration in the range of $0 \sim 50$ 515 nM according to the solar radiation (Figure S7). With this approach, the average of acetaldehyde seawater concentration estimated 516 by the model is 13.4 nM, a reasonable level compared to predicted level by Wang et al. (2019).

517 After adding the oceanic source of acetaldehyde, the model estimation was significantly improved (Figure 7). As the oceanic source 518 in the model is scaled according to the solar radiation, the measurement-to-model ratios were more strongly reduced during the 519 day compared to the night. With oceanic emission included, the model underestimation was less significant, within a factor of 3 520 during the day and 4 during the night over the Mediterranean Sea, Red Sea and Gulf of Aden. The most significant improvement 521 was identified over the Red Sea North. As shown in Figure 8, the model had much better agreement with the measurement after 522 adding the oceanic source. The scatter plots for other regions can be found in Figure S8. Over the Arabian Sea, the model 523 significantly overestimated acetaldehyde mixing ratios, indicating the input sea water concentration of acetaldehyde might be too 524 high. The SML layer starts to be effectively destroyed by the wave breaking when the wind speed exceeds than 8 m s⁻¹ (Gantt et 525 al., 2011). As the average wind speed over the Arabian Sea was the highest among the cruised areas (8.1 \pm 2.4 m s⁻¹, Figure S1), 526 less contribution from the CDOM photo degradation to acetaldehyde in the surface sea water would be expected. For the Suez 527 region, due to the limited model resolution $(1.1^{\circ} \times 1.1^{\circ})$, little sea water was identified in the model, leading to negligible influence 528 from the oceanic source.

529 Model underestimation of acetaldehyde especially over the Suez, Red Sea and Arabian Gulf is also likely to be related to the coarse 530 model resolution ($\sim 1.1^{\circ} \times 1.1^{\circ}$) (Fischer et al., 2015). Where model grid points contain areas of land the higher and more variable 531 terrestrial boundary layer height impacts the model prediction whereas the measurements may only by influenced by a shallower 532 and more stable marine boundary layer.



Figure 7. Acetaldehyde measurement to model ratios without the oceanic source (white boxes) and with the oceanic source (blue boxes) in the model during (a) daytime and (b) nighttime in different regions. The boxes represent 25% to 75% of the data set with the central line and square indicating the median and mean values, respectively. The whiskers show data from 10% to 90%.

537 The red dashed lines represent the 1:1 ratio.



539 Figure 8. Observed and simulated mixing ratios of acetaldehyde over the Red Sea North without oceanic emission (left) and with 540 oceanic emission (right). The data points are separated into day- and nighttime according to solar radiation.

541 3.4.3

538

Anthropogenic primary sources

542 Over the Arabian Gulf and Suez, the intensive photochemical production of carbonyls is apparent. Bourtsoukidis et al. (2020) 543 compared measured hydrocarbons (ethane, propane, and butane) with the results from model simulations (the same model used in 544 this study with the newly discovered deep water source implemented). The model was able to reproduce the measurement over 545 most regions expect for some significant model underestimations in Suez and Arabian Gulf, in which local and small-scale 546 emissions were difficult for the model to capture. Therefore, an underestimation of the precursor hydrocarbons, as well as those 547 large alkanes, alkenes and cyclic hydrocarbons which were not measured (> C8) or included in the model (> C5) could be a reason 548 for the model underestimation of acetaldehyde especially in polluted regions. In addition, as mentioned in the previous case studies, 549 high ozone mixing ratios were observed over the Arabian Gulf especially during the nighttime. Ethene and propene were found to 550 be significantly underestimated during the nighttime high ozone period by a factor over 10 (Figure S9), which indicates that the 551 nighttime ozonolysis of alkenes could be another important source for acetaldehyde, formaldehyde and other carbonyls (Atkinson 552 et al., 1995; Altshuller, 1993) in the Arabian Gulf.

553 Acetaldehyde, an oxygenated VOC, is not generally considered as an important primary emission from oil and gas field but instead 554 a photochemical product of hydrocarbon oxidation (Yuan et al., 2014;Koss et al., 2015;Koss et al., 2017). In contrast, primary 555 sources of formaldehyde from oil and gas production processes including both combustion and non-combustion process have been 556 ascertained (Vaught, 1991). Le Baron and Stoeckenius (2015) concluded in their report of the Uinta Basin winter ozone study that 557 besides formaldehyde, the other carbonyls were poorly understood in terms of their primary sources. Acetaldehyde and other 558 carbonyls (aldehydes and ketones) have been reported as primary emissions from fossil fuel combustion including ship emissions 559 (Reda et al., 2014;Xiao et al., 2018;Huang et al., 2018) and vehicle emissions (Nogueira et al., 2014;Erickson et al., 2014;Dong et 560 al., 2014). A possible explanation for the measurement-model discrepancy is that the active petroleum industry located in the 561 Arabian Gulf and intensive marine transportation in Suez are primary sources of acetaldehyde and other carbonyls which were not 562 well constrained in the model. The Suez region, where the largest acetaldehyde discrepancy was identified, had a significant 563 influence from biomass burning (see section 3.2.2). Biomass burning emissions are notoriously difficult to model as they are highly 564 variable both in time and space. In this study, the model failed to reproduce the acetonitrile level with a range of only 40-50 ppt 565 rather than 100-550 ppt measured over Suez. Thus, besides the possibility of seawater emission from the Gulf of Suez and the Suez 566 Canal, the underestimated biomass burning source in the model over Suez, will lead to an underestimation of acetaldehyde as well 567 as other carbonyl compounds in this region.

568 **3.4.4** Other possible secondary formation pathways

569 Although the model estimation was generally improved with the addition of an oceanic source, the model to measured ratios still 570 varied over a wide range. As mentioned above, photodegradation of CDOM on the surface of seawater is a known source for 571 acetaldehyde although some studies focusing on real sea water samples did not observe clear diel cycles of seawater acetaldehyde 572 (Beale et al., 2013; Yang et al., 2014). Fast microbial oxidation could be a reason (Dixon et al., 2013) while other non-light driven 573 sources of acetaldehyde could be an alternative explanation. In a recent study, Zhou et al. (2014) reported enhanced gas-phase 574 carbonyl compounds including acetaldehyde during a laboratory experiment of ozone reacting with SML samples, indicating 575 acetaldehyde could also be produced under non-light driven heterogeneous oxidation. Wang et al. (2019) ventured a hypothetical 576 source that organic aerosol can be an extra source for unattributed acetaldehyde in the free troposphere through light-driven 577 production and ozonolysis. However, since the yield of acetaldehyde from such reactions is unknown, large uncertainties remain. 578 Previous studies have shown that the organic matter fraction was highest in smaller sea spray aerosols and that the aerosols contain 579 both saturated and unsaturated fatty acids originating from the seawater surface (i.e. SML) (Mochida et al., 2002;Cochran et al., 580 2016). Thus, for the AQABA campaign, both photodegradation and heterogeneous oxidation could occur on the surface of sea 581 spray and pollution associated aerosols, even over remote open ocean therefore being an extra source of acetaldehyde and other 582 carbonyl compounds.

583 Another acetaldehyde formation pathway reported is gas-phase photolysis of pyruvic acid (Eger et al., 2019b;Reed Harris et al., 584 2016), a compound mainly of biogenic origin. Pyruvic acid has been also observed in seawater (Kieber and Mopper, 1987;Zhou 585 and Mopper, 1997) and was found up to 50 nM in the surface water of easter pacific Ocean (Steinberg and Bada, 1984), while 586 acetaldehyde was not the major product of aqueous-phase photolysis of pyruvic acid (Griffith et al., 2013). Zhou and Mopper 587 (1997) pointed out that the net exchange direction for pyruvic acid is expected to be from the air to the sea due to high solubility, 588 with Henry's law constant of 3.1×10^3 mol m⁻³ Pa⁻¹ (Sander, 2015). Moreover, partitioning to aerosols could be an important sink 589 for pyruvic acid (Reed et al., 2014; Griffith et al., 2013) : an increasing concentration trend of pyruvic acid was observed in marine 590 aerosols over western North Pacific Ocean (Boreddy et al., 2017). Therefore, due to limited terrestrial biogenic sources of pyruvic 591 acid for AQABA campaign, gas-phase level of pyruvic acid was expected to be low. Limited studies reported pyruvic acid level 592 in marine boundary layer, Baboukas et al. (2000) measured 1.1 ± 1.0 ppt of pyruvic acid above Atlantic Ocean. Pyruvic acid was 593 measured by Jardine et al. (2010) using a PTR-MS at m/z 89 in a forested environment. For the AQABA PTR-ToF-MS data set, 594 enhanced signals were observed at m/z 89.024 with the mean mixing ratio of 35-110 ppt over different regions (Table S4), which 595 is much more abundant than reported pyruvic acid levels by Baboukas et al. (2000). This might be due to the uncertainty associated 596 with the theoretical methods of quantification used here or the presence of isomeric compounds on that mass, since pyruvic acid 597 was not calibrated with the standard. Even if we assume the m/z 89.024 to be entirely pyruvic acid, with 60% yield of acetaldehyde 598 via photolysis (IUPAC, 2019), it gave maximum 13 ppt of acetaldehyde over Arabian Gulf, 5-9 ppt over other regions, which were 599 only 0.8% - 6% of the mean mixing ratios (Table S4). Detailed information of the calculation can be found in the Supporting 600 Information. Therefore, we conclude that the contribution from the photolysis of pyruvic acid is not an important source for the 601 unattributed acetaldehyde during the AQABA campaign.

602 4 Summary and Conclusion

603 Observations of carbonyl compounds around the Arabian Peninsula were investigated in terms of mixing ratios abundance over 604 different areas. Aliphatic carbonyl compounds were generally more abundant than the unsaturated and aromatic carbonyl 605 compounds, and were dominated by low-molecular-weight compounds (carbon number less than five). Aliphatic carbonyl 606 compounds were found at the highest mixing ratios over the Arabian Gulf followed by the Suez region, while the lowest mixing 607 ratios were observed over the Arabian Sea and the Gulf of Aden. Over the Mediterranean Sea, aliphatic carbonyls were low except 608 for acetone that was much higher compared to the levels observed over clean remote areas (i.e. Arabian Sea). The atmospheric 609 composition over the Red Sea showed obvious differences between the northern and the southern part, with higher mixing ratios 610 in the north. Similar region-dependent distributions were observed for unsaturated and aromatic carbonyls. Generally, the mixing 611 ratios of aromatic carbonyl compounds decreased as the carbon number increased. Particularly over the Suez region, benzaldehyde 612 (C7 aromatic carbonyls) was much more abundant than other aromatic carbonyls, indicating direct sources as well as abundant 613 oxidation precursors. For unsaturated carbonyl compounds, C5 and C6 carbonyl compounds dominated the mixing ratio 614 distribution, while the air chemistry highly depends on the chemical structure assignment of those masses.

Further case studies showed that the carbonyl compounds were highly correlated to the high ozone levels during daytime over the Arabian Gulf while the air chemistry in Suez region was strongly influenced by regional biomass burning. Due to the unexpectedly high loading of m/z 69 (usually assigned as isoprene) observed in highly polluted regions, we further identified the correlations between m/z 69 and other fragmentation masses of cycloalkanes according to previous studies conducted in oil and gas regions (Warneke et al., 2014;Yuan et al., 2014;Koss et al., 2017). The high correlations among fragments implied the existence of cycloalkanes in the polluted regions, which could be further oxidized to unsaturated carbonyl compounds (cyclic ketones or aldehydes).

622 As acetaldehyde was identified as having important additional sources, we further compared the measurements of major carbonyl 623 species (acetaldehyde, acetone and MEK) with a comprehensive global atmospheric chemistry model (EMAC). Acetaldehyde was 624 found to have the highest discrepancy between the observations and model simulations, with the simulated values to be lower up 625 to a factor of 10. By adding an oceanic source of acetaldehyde produced via light-driven photodegradation of CDOM in the 626 seawater, the model estimation improved significantly, especially over the Red Sea North. With the oceanic source added, modelled 627 acetaldehyde became slightly overestimated in clean regions, suggesting that the emission rate employed represents an upper limit. 628 The results indicate that the ocean plays an important role in the atmospheric acetaldehyde budget, under both clean and polluted 629 conditions. The underestimated acetaldehyde in the model is significant as it will influence the atmospheric budget of e.g. PAN. 630 As shown in Figure 1, multiple sources and formation pathways need to be considered to better understand the atmospheric budget 631 of acetaldehyde. Additional laboratory experiments and field measurements are necessary in order to verify all possible 632 atmospheric formation mechanisms and to improve model simulations.

633 Data availability.

634 The data used in this study are available to all scientists agreeing the AQABA protocol to at: 635 http://doi.org/10.5281/zenodo.3974228 (Wang et al., 2020).

636 Author contributions.

AE and CS performed PTR-ToF-MS measurement and preliminary data processing. NW conducted data analysis and drafted the article. AP performed EMAC model simulation. EB and LE are responsible for NMHC measurements and data. DD, BH and HF provided formaldehyde data. Ozone and actinic flux data were contributed by JS and JNC. Methane and carbon monoxide data were provided by JP. JL designed and realized the campaign. JW supervised the study. All authors contributed to editing the draft and approved the submitted version.

642 Competing interest.

643 The authors declare that they have no conflict of interest.

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