The driving factors of new particle formation and growth in the polluted boundary layer

Mao Xiao¹, Christopher R. Hoyle^{1,2}, Lubna Dada³, Dominik Stolzenburg⁴, Andreas Kürten⁵, Mingyi Wang⁶, Houssni Lamkaddam¹, Olga Garmash³, Bernhard Mentler⁷; Ugo Molteni¹, Andrea Baccarini¹,

- Mario Simon⁵, Xu-Cheng He³, Katrianne Lehtipalo^{3,8}, Lauri R. Ahonen³, Rima Baalbaki³, Paulus S. Bauer⁴, Lisa Beck³, David Bell¹, Federico Bianchi³, Sophia Brilke⁴, Dexian Chen⁶, Randall Chiu⁹, António Dias¹⁰, Jonathan Duplissy^{3,11}, Henning Finkenzeller⁹, Hamish Gordon¹², Victoria Hofbauer⁶, Changhyuk Kim^{13,14}, Theodore K. Koenig⁹, Janne Lampilahti³, Chuan Ping Lee¹, Zijun Li¹⁵, Huajun Mai¹³, Vladimir Makhmutov¹⁶, Hanna E. Manninen¹⁷, Ruby Marten¹, Serge Mathot¹⁷, Roy L.
- Mauldin^{18,19}, Wei Nie²⁰, Antti Onnela¹⁷, Eva Partoll⁷, Tuukka Petäjä³, Joschka Pfeifer^{5,17}, Veronika Pospisilova¹, Lauriane L. J. Quéléver³, Matti Rissanen^{3‡}, Siegfried Schobesberger¹⁵, Simone Schuchmann¹⁷, Yuri Stozhkov¹⁶, Christian Tauber⁴, Yee Jun Tham³, António Tomé²¹, Miguel Vazquez-Pufleau⁴, Andrea C. Wagner^{5,9‡‡}, Robert Wagner³, Yonghong Wang³, Lena Weitz⁵, Daniela Wimmer^{3,4}, Yusheng Wu³, Chao Yan³, Penglin Ye^{6,22}, Qing Ye⁶, Qiaozhi Zha³, Xueqin Zhou⁵, Antonio Amorim¹⁰,
- 15 Ken Carslaw¹², Joachim Curtius⁵, Armin Hansel⁷, Rainer Volkamer^{9,19}, Paul M. Winkler⁴, Richard C. Flagan¹³, Markku Kulmala^{3,11,20,23}, Douglas R. Worsnop^{3,22}, Jasper Kirkby^{5,17}, Neil M. Donahue⁶, Urs Baltensperger^{1*}, Imad El Haddad^{1*} & Josef Dommen¹
 - ¹Laboratory of Atmospheric Chemistry, Paul Scherrer Institute, 5232 Villigen, Switzerland
 - ²Institute for Atmospheric and Climate Science, ETH Zurich, 8092 Zurich, Switzerland
- 3 Institute for Atmospheric and Earth System Research (INAR) / Physics, University of Helsinki, 00014 Helsinki, Finland Faculty of Physics, University of Vienna, 1090 Vienna, Austria
 - ⁵Institute for Atmospheric and Environmental Sciences, Goethe University Frankfurt, 60438 Frankfurt am Main, Germany ⁶Center for Atmospheric Particle Studies, Carnegie Mellon University, Pittsburgh, PA 15213, USA
 - ⁷Institute of Ion Physics and Applied Physics, University of Innsbruck, 6020 Innsbruck, Austria
- 25 ⁸Finnish Meteorological Institute, 00560 Helsinki, Finland
 - ⁹Department of Chemistry & CIRES, University of Colorado Boulder, Boulder, CO 80305, USA
 - ¹⁰CENTRA and FCUL, University of Lisbon, 1749-016 Lisbon, Portugal
 - ¹¹Helsinki Institute of Physics, University of Helsinki, 00014 Helsinki, Finland
 - ¹²University of Leeds, LS2 9JT Leeds, United Kingdom
- 30 13Division of Chemistry and Chemical Engineering, California Institute of Technology, Pasadena, CA 91125, USA
 - ¹⁴School of Civil and Environmental Engineering, Pusan National University, 46241 Busan, Republic of Korea
 - ¹⁵Department of Applied Physics, University of Eastern Finland, 70211 Kuopio, Finland
 - ¹⁶P.N. Lebedev Physical Institute of the Russian Academy of Sciences, 119991 Moscow, Russian Federation ¹⁷CERN, CH-1211, Geneva, Switzerland
- 35 ¹⁸Department of Chemistry, Carnegie Mellon University, Pittsburgh, PA 15213, USA
 - ¹⁹Department of Oceanic and Atmospheric Sciences, University of Colorado Boulder, Boulder, CO 80309, USA
 - ²⁰Joint International Research Laboratory of Atmospheric and Earth System Sciences, School of Atmospheric Sciences, Nanjing University, Nanjing, Jiangsu Province, China
 - ²¹IDL-Universidade da Beira Interior, Covilhã, Portugal
- 40 ²²Aerodyne Research Inc., Billerica, MA 01821-3976, USA
 - ²³Aerosol and Haze Laboratory, Beijing Advanced Innovation Center for Soft Matter Science and Engineering, Beijing University of Chemical Technology, Beijing, China.
 - [‡]now at: Aerosol Physics Laboratory, Physics Unit, Faculty of Engineering and Natural Sciences, Tampere University, Tampere, Finland
- 45 **now at: Department of Chemistry & CIRES, University of Colorado Boulder, Boulder, CO 80305, USA

Abstract.

New-particle formation (NPF) is a significant source of atmospheric particles, affecting climate and air quality. Understanding the mechanisms involved in urban aerosols is important to develop effective mitigation strategies. However, NPF rates reported in the polluted boundary layer span more than four orders of magnitude and the reasons behind this variability subject of intense scientific debate. Multiple atmospheric vapours have been postulated to participate in NPF, including sulfuric acid, ammonia, amines and organics, but their relative roles remain unclear. We investigated NPF in the CLOUD chamber using mixtures of anthropogenic vapours that simulate polluted boundary layer conditions. We demonstrate that NPF in polluted environments are largely driven by the formation of sulfuric acid-base clusters, stabilized by the presence of amines, high ammonia concentrations and lower temperatures. Aromatic oxidation products, despite their extremely low volatility, play a minor role in NPF in the chosen urban environment but can be important for particle growth and hence for the survival of newly formed particles. Our measurements quantitatively account for NPF in highly diverse urban environments and explain its large observed variability. Such quantitative information obtained under controlled laboratory conditions will help the interpretation of future ambient observations of NPF rates in polluted atmospheres.

1 Introduction

New-particle formation (NPF) is an important atmospheric phenomenon, affecting both climate (Dunne et al., 2016) and air quality (Guo et al., 2014). Extremely high NPF rates are frequently observed in the polluted boundary layer, although current understanding suggests that newly formed particles should be rapidly scavenged by the high concentration of preexisting aerosols (Kulmala et al., 2017). Different vapours have been postulated to participate in NPF, including sulfuric acid, ammonia (Kirkby et al., 2011; Kürten et al., 2016), amines (Almeida et al., 2013), and organics (Kirkby et al., 2016; Lehtipalo et al., 2018; Riccobono et al., 2014).

The high NPF rates, believed to drive haze events in China (Guo et al., 2014), have been associated with the nucleation of sulfuric acid (H₂SO₄) in the presence of amines (Yao et al., 2018). In contrast, at other urban locations (Kuang et al., 2008), reported NPF rates are several orders of magnitude lower at similar H₂SO₄ concentrations, despite high levels of condensable species able to grow newly formed particles. Cai et al. (2020) attribute NPF in Beijing to H₂SO₄-amine cluster formation, which is modulated by coagulation scavenging. In Barcelona, Brean et al. (2020) linked nucleation to sulfuric acid clustering with both highly oxygenated organic molecules (HOMs) and bases, while Guo et al. (2020) argued photooxidation of organics from vehicular exhaust is responsible for the formation of ultrafine particles in Beijing. These conflicting observations and

interpretations highlight the need to better understand the role of the different vapours and environmental parameters and to quantify their relative contribution in new-particle formation and growth in different polluted locations.

The high NPF rates, believed to drive haze events in China (Guo et al., 2014), have been associated with the nucleation of sulfuric acid (H₂SO₄) in the presence of amines (Yao et al., 2018). In contrast, at other urban locations (Kuang et al., 2008), reported NPF rates are several orders of magnitude lower at similar H₂SO₄-concentrations, despite high levels of condensable species able to grow newly formed particles. These conflicting observations are subject of intense scientific debate of late (Brean et al., 2020; Cai et al., 2020; Guo et al., 2020) and highlight the need to better understand the role of the different vapours and environmental parameters and quantify their relative contribution in new-particle formation and growth in different polluted locations.

Here we determine the parameters controlling particle formation and growth under polluted boundary layer conditions in the CERN CLOUD chamber (Cosmics Leaving OUtdoor Droplets (Kirkby et al., 2011)). We investigated a complex mixture of H₂SO₄, ammonia, dimethylamine (DMA), NO_x, ozone, water and several anthropogenic volatile organic compounds (AVOCs: naphthalene (NAPH), 1,2,4-trimethylbenzene (TMB) and toluene (TOL)), at different temperatures. Organic vapours from the oxidation of anthropogenic precursors are expected to contribute to formation and growth but their role is not yet quantified. The experiments cover low and moderately polluted rural environments as well as highly polluted urban situations.

2 Results

100

105

2.1 Nucleation rates

NPF rates in the atmosphere often exhibit a clear correlation with H_2SO_4 (Kuang et al., 2008; Paasonen et al., 2010; Yao et al., 2018), but chamber experiments show that H_2SO_4 and water alone cannot explain boundary layer NPF events (Kirkby et al., 2011). Figure 1 presents our measured particle formation rates (at 1.7 nm diameter, referred to as $J_{1.7}$) as a function of H_2SO_4 , in the presence of various concentrations of ammonia, DMA, oxidised anthropogenic organic vapours, and NO_x , on top of ambient urban observations (sub-2 nm J). Though they are highly correlated with the H_2SO_4 concentration for otherwise fixed conditions, the CLOUD formation rates span more than five orders of magnitude for the same H_2SO_4 concentration at different conditions, similar to the ambient observations, both in their magnitude and dependence on H_2SO_4 (Kuang et al., 2008; Paasonen et al., 2010; Yao et al., 2018).

Most of the variation in the CLOUD measurements is driven by DMA. Despite much lower ambient concentrations compared to ammonia, amines can be a key driver of boundary layer NPF (Almeida et al., 2013; Yao et al., 2018). At 293 K, the addition of 4 pptv DMA to an NH₃ / H₂SO₄ mixture increases the particle formation rates by two to three orders of magnitude (magenta squares compared to open red triangles). At these relatively high temperatures, NPF rates do not reach the kinetic limit of H₂SO₄ nucleation (solid cyan line, Fig. 1), suggesting that higher amine levels could increase NPF rates further by 1 - 2 orders of magnitude.

In the absence of amines, $J_{1.7}$ is extremely sensitive to ammonia concentrations (Fig. 2A). Even with 1 - 2 ppbv NH₃ at 293 K (Fig. 1), atmospherically relevant H₂SO₄ concentrations result in NPF rates only near the lower end of those observed in the polluted boundary layer. These rates can increase by a factor of 100 when NH₃ rises to 10 ppbv (Fig. 2A), a level that is frequently found in polluted cities (Elser et al., 2018; Guo et al., 2017). Our new measurements at higher ammonia levels extend the previous parameterisation of Dunne et al. (2016). and reveal a $J_{1.7} \propto [\text{NH}_3]^2$ dependence within the atmospherically relevant range of [NH₃], in agreement with kinetic nucleation modelling (see Materials and Methods) based on thermodynamic data (Kürten, 2019) derived from CLOUD measurements and quantum chemical calculations.

NPF rates strongly increase with decreasing temperature (Fig. 1). The formation rate of H₂SO₄-NH₃ particles increases by two orders of magnitude at 278 K (blue line) compared to 293 K (red line). A similar increase in the formation rates of H₂SO₄-DMA particles is observed with the same temperature decrease (cyan line at 278 K vs. magenta line at 293 K). Accordingly, NPF at 278 K proceeds close to kinetic limit with only 4 pptv of DMA, more than 4 orders of magnitude faster than our baseline experiments (H₂SO₄ with 1 ppbv NH₃ at 293K, red dashed line). This agrees with J_{1.7} values extrapolated by Kürten et al. (2018) from measurements at 3.2 nm in Almeida et al. (2013). We modelled the temperature dependence of J_{1.7} for both H₂SO₄-NH₃ and H₂SO₄-DMA systems (Fig. S1) and find that H₂SO₄-base NPF rates can be reasonably well simulated at both temperatures, based on experimental and quantum-chemically-calculated thermodynamic data (Kürten, 2019; Myllys et al., 2019).

125

130

135

140

To investigate the effect of organic vapours, we initiated the photo-oxidation of a mixture of AVOCs (naphthalene, 1,2,4-trimethylbenzene and toluene), in the presence of SO₂, NH₃ and 0.1 - 1.5 ppbv NO. *J*_{1.7} increased by a factor of 2 - 100 compared to experiments without AVOCs (red filled triangles in Fig. 1A). Enhancements in NPF rates are significant in the absence of DMA, for the H₂SO₄-NH₃ system at 293 K (red dashed line), where NPF rates are lowest. Under these conditions, variations in the enhancement factors are explained by the NO_x levels and the amounts of oxidised organics (OxOrg) (Fig. 2B-C, Fig. S2, Fig. S3). However, *J*_{1.7} strongly decreases with increasing NO, since NO suppresses both autoxidation and the formation of low-volatility dimers. Compared with our baseline experiments, organics enhance *J*_{1.7} by a factor of 20 at 0.2 ppbv NO, but only by a factor of 2 at 1.2 ppbv NO, i.e., *J*_{1.7} is almost the same as for H₂SO₄ nucleating with NH₃ alone (Fig. 2D). It should also be noted that the HOMs dimer formation could be higher in chamber simulations like CLOUD than in the ambient atmosphere where extra HO₂ sources such as photolysis of aldehydes or OH reacting with CO or small oxygenated hydrocarbons are present (see SI for more details). In the presence of amines, the contribution of organics to NPF is marginal, since the inorganic nucleation rate is overwhelming. This is even more pronounced at higher NO mixing ratios, typical of urban atmospheres, where the influence of organics is further diminished.

Fig. 3 recapitulates the changes in NPF rates (J_{L7}) resulting from different factors (besides NH₃ and NO concentrations discussed in Fig. 2A and 2D) under fixed H₂SO₄ concentration. Lower temperature stabilizes nucleating clusters and enhances NPF rates under all conditions. While the addition of both DMA and organics leads to an enhancement in J_{L7} , the effect of only 4 pptv DMA is much more significant. Further addition of organics when DMA is already present will only marginally affect nucleation rates.

2.2 Growth of newly formed particles

145

150

155

160

165

170

In view of their high diffusivity, the growth of newly-formed particles up to 3 nm and above is crucial for their survival against scavenging loss. Condensation of H₂SO₄ and oxidised organics, as well as cluster coagulation contribute to growth (Lehtipalo et al., 2018). H₂SO₄ grows sub-3 nm particle by (1.94±0.13)× 10⁻⁷ nm h⁻¹ cm³, insufficient for growth rate (GR) measured at urban location. Such GR are consistent with the parameterisation in Lehtipalo et al. (2018) and model predictions including van-der-Waals interactions (Stolzenburg et al., 2020), while higher than estimates that assume the collision of hard spheres at the kinetic limit (Nieminen et al., 2010). DMA, low temperatures, and high NH₃ concentration (Lehtipalo et al., 2016) enhance molecular cluster concentration. When nucleation approaches the kinetic limit, cluster self-coagulation dominates the growth-cluster coagulation dominates the growth. As stable H₂SO₄–DMA clusters, formed to a great extent at 278 K, are lost less rapidly to pre-existing aerosols than are individual molecules (monomers), these clusters increase the condensable H₂SO₄ species reservoir and, thereby, the H₂SO₄ contribution to growth relative to organic vapours as seen in Fig. 4B (cyan squares). This increases the effective concentration of H₂SO₄ beyond the measured bare monomers, especially in the presence of high condensation sinks found in polluted environments.

While NPF is dominated by the formation of H₂SO₄ clusters, anthropogenic organic vapours strongly contribute to growth, and are required to explain the GR measured at urban locations. The multicomponent system used in the simulations this study does reproduce the range of GR found in urban atmospheres (4-12 nm h⁻¹ for 10⁷ cm⁻³ H₂SO₄). The measured GR are consistent with a 0.5 - 3% yield of OxOrg that are able to condense onto particles in the sub-3nm size range (Fig. S4).

Intense NPF events have been observed under highly polluted conditions, despite the high loss rates of nucleated clusters to pre-existing particles. Here, we investigated the survival probability at condensation sinks up to 0.02 s^{-1} . We define the survival probability as the ratio $J_{2.5}/J_{1.7}$ (see Fig. S5 for $J_6/J_{1.7}$) and display $J_{2.5}/J_{1.7}$ against the condensation sink, CS_6 , in Fig. 5, where CS_6 only includes particles larger than 6 nm that are typically reported from scanning mobility particle sizer measurements. For the case of H_2SO_4 -N H_3 nucleation, the survival probability decreases with increasing CS_6 , and increases with increasing GR. When NPF rates approach the kinetic limit in the presence of DMA, as, for example, in Shanghai (Yao et al., 2018), collisions between the abundant nucleating clusters reduce the survival probability of sub-3 nm particles by an order of magnitude, but simultaneously increase the growth rate of those that do survive. The survival probability only becomes dependent on the CS_6 when CS_6 is greater than the loss rate due to cluster-cluster collisions. For 10^7 cm⁻³ H_2SO_4 , the measured survival probability of newly formed particles remains virtually independent of CS_6 even for values as high as 0.01- 0.02 s^{-1} (Figure 5, Fig. S5), consistent with model simulations (black line). This is much higher than previously estimated (Kulmala et al., 2017; Yao et al., 2018) when van-der-Waals interactions were not considered (Fig. S6), and explain the low dependence of particle formation on CS observed even in highly polluted urban environments. Above 0.02 s^{-1} , CS_6 becomes the dominant sink of clusters, sharply decreasing their survival probability.

Figure 6 shows a mass defect plot of the nucleating species during an experiment with H₂SO₄, NH₃, DMA, NO_x, and oxidised anthropogenic organics at 278 K, representative of recent ambient measurements in Shanghai (Yao et al., 2018). Similar

observations were also presented for the case of Beijing (Cai et al., 2021). In the same way as in these ambient observations, pure H₂SO₄ clusters are seen up to tetramers. Larger acid clusters contain the stronger base, DMA. While the precursor composition in the ambient atmosphere is indeed much more complex than in the chamber, we demonstrate that the features of HOMs in both cases are very similar, representative of high NO_x conditions. This includes the high abundance of HOMs with a molecular weight between 200-400 Th, no significant monomer/dimer patterns, and a high contribution of N-containing HOMs. These observations indicate that, under similar conditions, chamber simulations reproduce ambient observations.

2.3 *J* vs GRs

185

190

195

200

If the presence of amines (Cai et al., 2021; Yao et al., 2018) or organics (Brean et al., 2020; Guo et al., 2020) drives urban NPF has been an intense debate in recent atmospheric observations. While both organics and amines can enhance H₂SO₄-NH₃ nucleation to different extents (Fig. 3), they do so by different mechanisms. Amines stabilize the nucleating clusters while organics contribute to higher concentrations of condensing material. We compare chamber simulations and atmospheric observations in Fig. 7 using sub-3 nm GR as a proxy for total condensable vapour concentrations, as most field observations in urban atmospheres do not report the nucleating species at a molecular level. For H_2SO_4 -base nucleation, $J_{1,7}$ and sub-3 nm GR are tightly correlated through their dependence on H₂SO₄ concentration. At a fixed GR, J values are enhanced by lower temperatures and by DMA compared with ammonia. In line with this, measurements by Yao et al. (2018) in Shanghai fall on the H₂SO₄ kinetic limit nucleation line. Shanghai NPF shows no seasonal variation, implying that DMA is present year round and that concentrations during the warm seasons are higher than 4 pptv. NPF during wintertime in Nanjing (Yu et al., 2016) also falls on the H₂SO₄ kinetic limit nucleation line, and so is likely driven by amines as well. The CLOUD measurements of NPF at 278 K with H₂SO₄, 1-2 ppbv ammonia and organics compare well with atmospheric observations in several polluted environments during cold seasons (Beijing winter; Jayaratne et al., 2017) and Yangtze River Delta region (Dai et al., 2017)) implying low amine concentrations (< 4 pptv) in these regions in winter. Atmospheric observations in warm environments lie well above the J vs GR line for 1 to 2 ppbv ammonia nucleating with H₂SO₄ at 293 K. Ammonia concentrations of ~10 ppbv could explain the observations from Madrid (Carnerero et al., 2018), Po Valley (Kontkanen et al., 2017) and Tecamac (Kuang et al., 2008), while unrealistically high ammonia concentrations (>100 ppbv) would be needed to explain Beijing spring (Cai et al., 2017) or Nanjing summer (Yu et al., 2016). Therefore, NPF in the latter two cases is likely due to amines at ppty levels, which would match the CLOUD measurements well (magenta squares in Fig. 7). Looking at the variability in NPF observed in Shanghai, Nanjing (summer and winter) and the Yangtze River Delta, we see evidence for strong regional variation in polluting vapours. This controlled laboratory exploration spans the relevant ambient conditions, and provides a detailed understanding and constraints of the governing chemical and physical processes of NPF and early growth. This provides the possibility to rationalize the available ambient measurements and a framework to plan future measurements in the best way.

205 3 Conclusions

210

215

230

235

Urban NPF rates are primarily driven by the ambient concentrations of H₂SO₄ and the presence of stabilising base vapours. The largest effect results from amines, if present, as even low pptv levels are sufficient for NPF rates to approach the kinetic limit for H₂SO₄ particles. Reducing amine emissions may therefore be an efficient way to reduce the particle number concentrations, although the particle mass concentrations will not be strongly affected since the available condensable vapours will condense onto pre-existing particles. Although ammonia is less efficient for stabilizing H₂SO₄ clusters, reducing the NH₃ concentrations can also bring strong benefits in conditions where amine concentration is limited, due to the quadratic dependence of NPF on NH₃. Lower temperatures also strongly stabilize clusters, suggesting that particle number concentrations may be affected by a warming climate. In spite of the high VOC emissions in cities, the contribution of organics to nucleation is likely to be limited by high co-emission of NO_x, although organic oxidation products remain important for particle growth. In contrast, in suburban environments, an enhanced organic contribution to NPF is expected because of lower NO_x and base concentrations. We conclude that in highly polluted environments such as Asian megacities with high ammonia and NO_x concentrations and, especially, in the presence of amines, H₂SO₄–base nucleation is the prime driver of new particle formation and is effective even in the presence of high condensation sinks near 0.02 s⁻¹.

4 Materials and Methods

The experiments are performed at CERN CLOUD chamber (Cosmics Leaving OUtdoor Droplets (Kirkby et al., 2011)) with a complex mixture of H₂SO₄, ammonia, dimethylamine (DMA), NO_x, ozone and several anthropogenic volatile organic compounds (AVOCs: naphthalene (NAPH), 1,2,4-trimethylbenzene (TMB) and toluene (TOL)), at 293K and 278 K, 60%. Organic vapours from the oxidation of anthropogenic precursors, besides their effect on human health (Daellenbach et al., 2020) are expected to contribute to formation and growth, similar to biogenic products (Lehtipalo et al., 2018), but their role is still largely unknown. The data were collected between October and November in 2016 and 2017

4.1 Chamber experiments

Experiments were performed at the CLOUD facility (Kirkby et al., 2011) at CERN during two intensive campaigns in 2016 (CLOUD11 campaign) and 2017 (CLOUD12 campaign). All experiments were carried out at atmospheric pressure with ultrapure air created from the evaporation of liquid nitrogen (Messer, 99.999%) and liquid oxygen (Messer, 99.999%), which were mixed in the gas volume ratio 79% and 21%, respectively. The O₃ was generated by exposing a small fraction of the air through a quartz tube surrounded by UVC lamps (wavelength <240 nm). Relative humidity was adjusted by passing the air through a Nafion® humidifier using ultrapure water (18 MΩ cm, Millipore Corporation). The required SO₂ mixing ratio was provided from a pressurized gas cylinder (CARBAGAS AG, 500 ppmv in N₂). The anthropogenic volatile organic compounds were added by flushing air through an evaporator containing naphthalene (Sigma Aldrich, 99%) or from gas bottles of 1,2,4-

trimethylbenzene (Messer) and toluene (Messer). With all the instruments connected to the CLOUD facility, the total sampling flow rate was 270 L min⁻¹, resulting in a dilution lifetime of 1.6 h in the chamber. All experiments presented here were performed in the presence of ions from galactic cosmic rays (GCR) at normal atmospheric levels, i.e. without ion removal using field cage electrodes, except for data specified as neutral experiments, which includes data without ions where the field cage electrodes were set to $\pm 30 \mathrm{kV}$.

240

245

250

255

260

265

In the CLOUD11 campaign, experiments were performed at 293±1.5 K (referred to as 293 K) and 57±3 % RH. Approximately 1 to 2 ppbv NH₃ were injected into the chamber from an NH₃ gas bottle (100 ppm in N₂, CARBAGAS AG). A set of UV sources including a 4 W KrF excimer UV laser (UVX) at 248 nm and four 200 W Hamamatsu Hg-Xe lamps (UVH) at wavelengths between 250 and 450 nm with adjustable power were used to photolyse 40 to 100 ppbv O₃ to produce 1 x 10⁶ to 1 x 10⁷ OH cm⁻³. Extra UVA at 385 nm was generated from a 400 W UVA LED saber (UVS3) to photolyse NO₂ to NO.

Three sets of experiments were performed during CLOUD11: (1) experiments to determine NPF rates from sulfuric acid and ammonia, (2) experiments to determine NPF rates from sulfuric acid and typical anthropogenic volatile organic compounds (AVOCs), i.e., naphthalene (NAPH), 1,2,4-trimethylbenzene (TMB) and toluene (TOL) with and without ammonia, and (3) experiments to determine NPF rates from sulfuric acid, ammonia and AVOC mixtures in the presence of NO_x. A typical experiment started by turning on UV sources to trigger oxidation of AVOCs and SO₂. Formation rates (*J*) and growth rates (GRs) were determined by measuring the dynamics of the particle number size distribution (see below).

In the CLOUD12 campaign, experiments were performed at 278±1 K (referred to as 278 K) or 294±1 K (referred to as 293 K) and 60±5 % RH. 1 to 2 ppbv NH₃ were injected into the chamber from an NH₃ gas bottle (100 ppm in N₂, Messer). 4 pptv dimethylamine (DMA, 1% in N₂, Messer) was injected in the last part of the campaign to study the effect of amines on NPF. In addition to O₃ photolysis, OH radicals were also produced by nitrous acid (HONO) photolysis in this campaign. The gaseous HONO was synthesised by continuous mixing of H₂SO₄ (Sigma Aldrich, 99%) with NaNO₂ (Sigma Aldrich, 99%) in a specially designed stainless-steel reactor. 2 to 20 ppbv HONO was injected into the CLOUD chamber and photolysed by UVS3 producing 1 x 10⁶ to 1 x 10⁷ OH (for details see section 4.5) together with O₃ photolysis by UVH and UVX as in CLOUD11. Three sets of experiments were performed during CLOUD12 at 293K and 278K: (1) experiments in the presence of sulfuric acid, ammonia and AVOC mixtures; (2) experiments in the presence of sulfuric acid and DMA, and (3) experiments with sulfuric acid, DMA and AVOC mixtures.

It is expected that precursor VOCs must have a large carbon skeleton to be able to form highly oxygenated molecules (HOMs) contributing to NPF. Aromatics comprise roughly 20-40% of VOC emissions in urban environments (Baker et al., 2008; Boynard et al., 2014; Li et al., 2019), while the largest fraction is from small alkanes and alkenes that are not relevant for our study. The AVOC mixtures used in this study represent the most common aromatic compounds in the urban boundary layer. Toluene is the most abundant aromatic compound. We have not included benzene, as it has lower emission rates than toluene, and is unlikely to contribute to NPF due its low reactivity. TMB was chosen as representative of more reactive C2 and C3 aromatics, which individually have a factor of 1-5 smaller concentrations than toluene. Naphthalene is representative of polycyclic aromatic compounds, and, potentially, of other unspeciated organics (e.g., long-chain alkanes (Bruns et al., 2016;

Jathar et al., 2014)) which are emitted at lower rates but have high potential of forming SOA (Chan et al., 2009; Li et al., 2016) and HOMs (Molteni et al., 2018). We tested different concentrations of the three aromatics with different mixing ratios to represent the wide range of HOM forming organics in urban environments listed in Table S1.

4.2 Particle Measurements

275

280

285

Several independent instruments measured particle concentrations and number size distributions in the chamber. Number concentrations of freshly formed particles at several different cut-off sizes between 1 and 3 nm were determined by a particle size magnifier (Vanhanen et al., 2011) (PSM, Airmodus Ltd.) using diethylene glycol as working fluid. The PSM was used in scanning mode and the particle counting was done by a condensation particle counter (CPC) downstream of the PSM. A butanol ultra-fine CPC (TSI 3776) with fixed cut-off size measured the total particle number concentration above ca. 2.5 nm. The particle size distribution between 1.8 and 8 nm was measured with a differential mobility analyzer trainDMA train (Stolzenburg et al., 2017) which consists of six differential mobility analysers and CPCs in parallel. For the two channels measuring sub-2.5 nm particles, a PSM was installed upstream of the particle counter. A commercial nano-SMPS-scanning mobility particle sizer (nano-SMPS TSI 3938) with a water-CPC (TSI 3788) was used to measure the size distribution between 4.6 and 60 nm, and a custom-built SMPS, consisting of a krypton source, a long differential mobility analyzerDMA and a CPC (TSI 3010) was used to cover the size distribution between 20 and 400 nm. A neutral cluster and air ion spectrometer (Mirme and Mirme, 2013) (NAIS, Airel Ltd.) measured ions with mobility diameters between 0.75 and 40 nm. A corona charger was periodically applied to detect particle size distributions in the size range of 2-40 nm using the NAIS.

4.3 Determination of new particle formation rates

Particle formation rates (J_{dp}) were calculated as in a previous study (Dada et al., 2020):

$$J_{dp} = \frac{dN_{dp}}{dt} + S_{dil} + S_{wall} + S_{coag} \tag{1}$$

where N_{dp} is the particle number concentration above a certain cut-off size (d_p), determined by the PSM for 1.7 nm, the CPC for 2.5 nm and the nano-SMPS for 4.6 nm. The cut-off size has an uncertainty of approximately 0.5 nm.

Dilution was corrected for by applying the relation $S_{dil} = k_{dil}N_{dp}$, where $k_{dil} = 1.72 \times 10^{-4} \text{ s}^{-1}$ accounts for a 270 L min⁻¹ continuous flow into the 26.1 m³ chamber.

Particle diffusion loss to the chamber wall was calculated as:

$$S_{wall} = \sum_{d'_p = d_p}^{d_{p,max}} N_{d'_p} k_{wall}(d'_p, T)$$
(2)

where wall loss rates were extrapolated from the sulfuric acid wall loss rate:

$$k_{wall}(d'_p, T) = 2.116 \times 10^{-3} \left(\frac{T}{278 \, K}\right)^{0.875} \left(\frac{0.82 \, nm}{d'_p}\right)$$
 (3)

295 Coagulation losses were calculated as:

$$S_{coag} = \sum_{d_{p,i}=d_p}^{d_{p,max}} \sum_{d_{p,j}=d_{p,i}}^{d_{p,max}} \delta_{ij} K_{ij} N_i N_j$$

$$\tag{4}$$

where K_{ij} is the coagulation coefficient for particles of size $d_{p,i}$ and $d_{p,j}$, and N_i are the particle concentrations in the size bins i and j. δ =1 for $i\neq j$ and δ =0.5 for i=j. Combined size distributions based on nano-SMPS and SMPS were used to calculate coagulation losses. For some experiments in CLOUD12 nano-SMPS measurements were not available. Instead, NAIS measurements in particle mode (i.e., with corona charger), cross validated with nano-SMPS measurements, were used to construct size distributions for particles below 20 nm.

The J values reported here are the median values of each experiment after reaching stable conditions. The uncertainty of J (as error bars in the figures) was calculated by an error propagation method, accounting for the statistical variation of J in the experiment and run-to-run repeatability of J in CLOUD, which is ca. 30%.

Growth rates were determined by the appearance time method (Lehtipalo et al., 2014, 2016) using the 50% appearance time of the DMA-train measurements from 1.8 nm to 3.2 nm.

4.4 Gas phase measurements

300

310

315

Gas monitors were used to measure ozone (O₃, Thermo Envrionmental Instruments TEI 49C), sulfur dioxide (SO₂, Thermo Fisher Scientific Inc. 42i-TLE) and nitric oxide (NO, ECO Physics, CLD 780TR). Nitrogen dioxide (NO₂) was measured by a cavity attenuated phase shift nitrogen dioxide monitor (CAPS NO₂, Aerodyne Research Inc.) and a custom-made cavity enhanced differential optical absorption spectroscopy (CE-DOAS) instrument. Relative humidity of the chamber was determined by dew point mirrors (EdgeTech).

Two HNO₃-chemical ionisation mass spectrometers (Jokinen et al., 2012) (HNO₃-CIMS) were connected to the chamber for gas phase measurements. One instrument was an atmospheric pressure interface time of flight mass spectrometer (Junninen et al., 2010) (APi-TOF, TOFWERK AG) equipped with a commercial inlet (Airmodus, Ltd) using NO₃⁻ as primary ions. Another HNO₃-CIMS with an ion precipitator in front of the inlet used a corona discharge as ion source and 50 ns reaction time. Sulfuric acid was measured by these HNO₃-CIMS and quantified following the same calibration and loss correcting procedures described in previous publications (Kirkby et al., 2016; Tröstl et al., 2016). The sulfuric acid concentrations measured by the two instruments agreed with each other within 20%. To avoid interference from charged ions in the chamber, the sulfuric acid concentrations from the HNO₃-CIMS with an ion precipitator were used in the study.

Ammonia concentrations were measured with cavity ring-down spectroscopy (G2103, Picarro, Inc) and an H₃O⁺ chemical ionisation mass spectrometer (H₃O⁺ CIMS) (Pfeifer et al., 2020). The latter was an APi-TOF (TOFWERK AG) coupled with a home-made crossflow ionization source using positively charged water clusters to detect ammonia in real time. Ammonia

concentrations were measured with an H₃O⁺ chemical ionisation mass spectrometer (H₃O⁺ CIMS) and a PICARRO NH₃ analyser. The H₃O⁺ CIMS was an APi-TOF equipped with a crossflow inlet using positively charged water clusters to detect ammonia in real time. Dimethylamine was also measured by the H₃O⁺ CIMS applying the calibration factor determined from ammonia. The measured concentration of dimethylamine (DMA) was in good agreement with estimations based on DMA injection rates. DMA time traces could also be measured by the HNO₃-CIMS (Simon et al., 2016), which was used for experiments when the DMA signal in the H₃O⁺ CIMS was influenced by fragments from organics.

A <u>custom-made</u> proton transfer time of flight mass spectrometer (PTR3) (Breitenlechner et al., 2017) with a core sampling inlet, which transfers sample air through a tripole reaction chamber operated at 80 mbar, was used to measure the TMB and NAPH concentration. A quadrupole proton transfer mass spectrometer (PTR-MS with quadrupole, IONICON) measured the toluene concentration in CLOUD11 and a <u>custom-made short TOF</u> proton transfer mass spectrometer was used in <u>CLOUD12and STOF</u> proton transfer mass spectrometer (PTRS) was used in <u>CLOUD12</u> to measure toluene, TMB and naphthalene. For some experiments in <u>CLOUD111</u>, the quadrupole PTR-MS measurements were not available, where we determined the toluene concentration from the injection rates (MFC settings).

4.5 Determination of OH and OxOrg concentrations

The OH concentration was derived from the sulfuric acid concentration as well as from the amount of TMB reacted. To determine OH from the sulfuric acid concentration, first order sulfuric acid change rates were fitted with a 15-min time window. A sulfuric acid production rate was then obtained by subtraction of sulfuric acid losses to the chamber wall and to particles.

340 The OH concentration was determined as:

$$[OH] = \frac{P_{SA}}{k_{SA}[SO_2]} \tag{5}$$

where k_{SA} =8.6 x 10⁻¹³ cm³ s⁻¹.

325

330

335

The OH concentration was also determined by fitting the TMB time series using:

$$\frac{d[TMB]}{dt} = k_{inj} - k_{dil} + k_{TMB}[OH][TMB] \tag{6}$$

where k_{inj} is the injection rate of TMB and k_{dil} is the chamber dilution rate. The reaction rate constant for OH and TMB (k_{TMB}) is $3.25 \times 10^{-11} \, \text{cm}^3 \, \text{s}^{-1}$. OH concentrations calculated from the two methods agreed within 20% at 293 K without DMA addition. For other conditions, OH concentration estimates from the two methods correlated with each other but the method using sulfuric acid concentration underestimated the OH concentration as the formation of sulfuric acid clusters was not taken into consideration. Therefore, OH concentrations from the sulfuric acid method were corrected by applying correction factors based on the OH estimation from the TMB time series. In the analysis here, the OH time series from the sulfuric acid concentration were used.

350 Oxidised organics (OxOrg) concentrations were estimated with

$$\frac{d[OxOrg]}{dt} = k_{VOC}[OH][VOC] - (k_{dil} + k_{wall} + CS)[OxOrg]$$
(7)

The reaction rate constants k_{VOC} are 2.6 x 10^{-11} cm³ s⁻¹ for NAPH, 3.25e x 10^{-11} cm³ s⁻¹ for TMB and 5.8 x 10^{-12} cm³ s⁻¹ for toluene. Dilution, wall loss and condensation to particles are considered by k_{dil} , k_{wall} and CS assuming oxidised organics will not evaporate after they condensed on a wall or particles, which is a reasonable assumption for compounds relevant for nucleation and early growth.

355 4.6 Nucleation Model

A kinetic model based on the general dynamic equation (Seinfeld and Pandis, 2016) was used in order to assess acid-base nucleation. The model considers neutral clusters containing up to 15 sulfuric acid molecules and 50 geometric size bins between clusters containing 15 sulfuric acid molecules and 300 nm and uses a constant sulfuric acid production rate:

$$\frac{dN_{A_{1}B_{0}}}{dt} = Production + \sum \gamma_{A,A_{k}B_{j}} N_{A_{k}B_{j}} + \gamma_{A,A_{2}B_{0}} N_{A_{2}B_{0}} - N_{A_{1}B_{0}} \left(k_{dil} + k_{wall,A_{i}B_{i}} + \sum_{j=1}^{\infty} K_{A_{i}B_{i},j} N_{j} \right)$$
(8)

Evaporation was considered for clusters containing less than 5 sulfuric acid molecules. Clusters containing more bases than 360 acids were ignored as they are highly unstable.

$$\frac{dN_{A_{i}B_{i}}}{dt} = \frac{1}{2} \sum_{\substack{A_{j}+A_{k}=A_{i} \\ B_{j}+B_{k}=B_{i}}} K_{j,k} N_{A_{j}B_{j}} N_{A_{k}B_{k}} + \gamma_{A,A_{i}+1B_{i}} N_{A_{i}+1,B_{i}} + \gamma_{B,A_{i}B_{i}+1} N_{A_{i}B_{i}+1}
- N_{A_{i}B_{i}} \left(k_{dil} + k_{wall,A_{i}B_{i}} + \sum_{j=1}^{\infty} K_{A_{i}B_{i},j} N_{j} \right) - (\gamma_{A,A_{i}B_{i}} + \gamma_{B,A_{i}B_{i}}) N_{A_{i}B_{i}}$$
(9)

Clusters containing 5 to 15 sulfuric acid molecules are considered as stable and were treated the same regardless the base number.

$$\frac{dN_i}{dt} = \frac{1}{2} \sum_{j=1}^{i-1} K_{j,i-j} N_j N_{i-j} - N_i \left(k_{dil} + k_{wall,i} + \sum_{j=1}^{\infty} K_{i,j} N_j \right)$$
(10)

Then 50 geometric sized bins were applied, and the particles formed by collision of two particles were distributed between the two nearest bins as given in Eq. (11) in order to link continuous growth to discrete collisions.

$$\frac{dN_i}{dt} = \frac{1}{2} \sum_{j+k \to i} K_{j,k} N_j N_k P_{j,k} + \frac{1}{2} \sum_{j+k \to i-1} K_{j,k} N_j N_k (1 - P_{j,k}) - N_i \left(k_{dil} + k_{wall,i} + \sum_{j=1}^{\infty} K_{i,j} N_j \right)$$
(11)

365 Collision rates were calculated as:

$$K_{i,j} = \left(\frac{3}{4\pi}\right)^{\frac{1}{6}} \left(\frac{6k_B T}{m_i} + \frac{6k_B T}{m_j}\right)^{\frac{1}{2}} \left(V_i^{\frac{1}{3}} + V_j^{\frac{1}{3}}\right)^2 E_{i,j}$$
(12)

$$P_{j,k} = \frac{V_{i+1} - V_j - V_k}{V_{i+1} - V_i} \tag{13}$$

Where $V_i < V_j + V_k < V_{i+1}$.

 $E_{i,j}$ is the collision enhancement factor due to the van-der-Waals force, which can be linked to the Hamaker constant (H_A=6.4×10⁻²⁰ J), as described by Chan and Mozurkewich (Chan and Mozurkewich, 2001).

$$E_{i,j} = 1 + \frac{\sqrt{A'/3}}{1 + 0.0151\sqrt{A'}} - 0.186 \ln(1 + A') - 0.0163 \ln^3(1 + A')$$
(14)

$$A' = \frac{4H_A d_{p_i} d_{p_j}}{kT(d_{p_i} + d_{p_j})}$$
(15)

where d_{pi} and d_{pj} are the diameters of the colliding particles. A density of 1.7 g cm⁻³ was used for (NH₄)₂SO₄ in the calculations. 70 Evaporation of clusters containing less than five sulfuric acid molecules was taken into consideration using updated Gibbs free energies reported by Kürten (2019) (Table 1 in corrigendum) based on previously published CLOUD results for ammonia sulfuric acid nucleation. For simplification, only evaporation of one ammonia ($\gamma_{B,AB}$) or one sulfuric acid ($\gamma_{A,AB}$) molecule was taken into consideration. All evaporation rates were set to zero to calculate nucleation at the kinetic limit. Clusters with five and more sulfuric acid molecules were considered to have an equal number of acids and bases.

375 Wall loss was calculated as

$$k_{wall} = C\sqrt{D} \tag{16}$$

where C is a chamber specific constant and depends on chamber dimensions and air mixing (fan speed). C=0.77 cm⁻¹ s^{-0.5} was used in this study based on a measured sulfuric acid wall loss rate of 0.002 s⁻¹. D is the diffusion constant calculated as

$$D = \frac{kTC_c}{3\pi\eta d_p} \tag{17}$$

where k, T, $C_{\rm C}$, η , $d_{\rm P}$ are the Boltzmann constant, temperature, Cunningham slip correction factor, gas viscosity and particle diameter, respectively.

$$C_c = 1 + Kn * (1.142 + 0.558 * exp(-0.999/Kn));$$
(18)

When discussing effects of condensation sink, condensation sinks were considered as being composed of varying concentrations of particles with a diameter of 100 nm.

Author Contribution: M.X., D.S., A.K., M.W., H.L., B.M., U.M., A.B., M.S., X.-C.H., K.L., L.A., R.B., D.C., A.D., J.Du., H.F., V.H., C.K., T.K., J.L. C.L., Z.L., H.M., V.M., H.E.M., R.L.M., A.O., E.P., T.P., J.P., V.P., L.Q., M.R., Sie.S., Sim.S., Y.S., Y.J.T., A.T., M.V., A.W., R.W., Yo.W., L.W., D.W., Yu. W., C.Y., P.Y., Y.Y., Q.Z., X.Z. R.F., M.K., J.K., J.Do. prepared

the CLOUD facility and measurement instruments. X.M., C.R.H., L.D., H.L., U.M., A.B., T.K., M.R., A.T., L.W., A.H., K.C., V.R., R.C.F., M.K., D.W., J.K., N.M.D., U.B., I-E.H., J.Do. planed the experiments. M.X., C.R.H., L.D., D.S., M.W., H.L., O.G. B.M., U.M., A.B., M.S., X.-C.H., K.L., L.A., R.B., P.S.B., L.B., D.B., F.B., S.B., R.C., A.D., J.Du., H.F., V.H., C.K., J.L., C.P.L., Z.L., V.M., R.M., R.L.M., W.N., E.P., J.P., V.P., M.R., Sim.S., Y.J.T., A.T., M.V., A.W., L.W., D.W., Yu.W., C.Y., P.Y., Q.Y., X.Z., A.A., P.M.W., I-E.H., J.Do. collected the data. X.M., C.R.H., L.D., D.S., A.K., M.W., H.L., O.G., B.M., U.M., M.S., H.F., C.K., R.L.M., A.C.W., L.W., P.Y., R.F., J.D. analyzed the data. X.M., C.R.H., D.S., M.W., H.L., O.G., U.M., A.B., F.B., H.G., C.K., T.P., M.R., Sie.S., P.Y., J.C., A.H., R.V., R.C.F., K.M., D.W., J.K., N.M.D., U.B., I-E.H., J.Do. contributed to the scientific discussion. M.X., C.R.H, H.L., T.P., R.F., N.M.D., U.B., I-E.H., J.Do. contributed to the writing of the manuscript.

395

Data availability: Data related to this article are available upon request to the corresponding authors. Supplement.

Competing interests: The authors declare that they have no conflict of interest.

400

405

410

415

Acknowledgements

We thank CERN for supporting CLOUD with technical and financial resources, and for providing a particle beam from the CERN Proton Synchrotron. We thank tofTools team for providing programs for mass spectrometry analysis. We thank P. Carrie, L.-P. De Menezes, F. Josa, I. Krasin, O.S. Maksumov, M.V. Philippov, R. Sitals, A. Wasem, K. Ivanova for their contributions to the experiment. Funding: This research has received funding from the EC Seventh Framework Programme and European Union's Horizon 2020 programme (Marie Skłodowska Curie ITNs no. 316662 "CLOUD-TRAIN" and no. 764991 "CLOUD-MOTION"), Horizon 2020 (Marie Skłodowska-Curie Grant "Nano-CAVa" 656994), Horizon 2020 MC-COFUND Grant (665779), ERC Advanced ('ATM-GP' grant no. 227463), ERC-Consolidator Grant (NANODYNAMITE 616075), ERC-Starting grant (COALA, grant no. 638703, QAPPA, grant no. 335478), the Swiss National Science Foundation (no. 200021 169090, 200020 172602, 20FI20 172622), the U.S. National Science Foundation (grants AGC1439551, AGS1447056, AGS1531284, AGS1801574, AGS1801897, AGS1649147, AGS1801280, AGS1602086, 1801329), Wallace Research Foundation, German Federal Ministry of Education and Research(01LK1222A CLOUD-12 and 01LK1601A CLOUD-16), the Portuguese Foundation for Science and Technology (project no. CERN/FIS-COM/0014/2017), the Presidium of the Russian Academy of Sciences ("High energy physics and neutrino astrophysics" 2015), the Presidium of the Russian Academy of Sciences (the Program "Physics of Fundamental Interactions" 2017-2020), Austrian Science Fund (FWF, project number J3951-N36), the Austrian Science Fund (FWF; project no. P27295-N20), NASA graduate fellowship (NASA-NNX16AP36H), Academy of Finland (project number 299574, 307331 and 310682), CERN/FIS-COM/0014/2017.

References

doi:10.1016/j.atmosenv.2007.09.007, 2008.

- Almeida, J., Schobesberger, S., Kürten, A., Ortega, I. K., Kupiainen-Määttä, O., Praplan, A. P., Adamov, A., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Donahue, N. M., Downard, A., Dunne, E., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Guida, R., Hakala, J., Hansel, A., Heinritzi, M., Henschel, H., Jokinen, T., Junninen, H., Kajos, M., Kangasluoma, J., Keskinen, H., Kupc, A., Kurtén, T., Kvashin, A. N., Laaksonen, A., Lehtipalo, K., Leiminger, M., Leppä, J., Loukonen, V., Makhmutov, V., Mathot, S., McGrath, M. J., Nieminen, T., Olenius, T., Onnela, A., Petäjä, T., Riccobono,
- F., Riipinen, I., Rissanen, M., Rondo, L., Ruuskanen, T., Santos, F. D., Sarnela, N., Schallhart, S., Schnitzhofer, R., Seinfeld, J. H., Simon, M., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A., Tröstl, J., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y., Virtanen, A., Vrtala, A., Wagner, P. E., Weingartner, E., Wex, H., Williamson, C., Wimmer, D., Ye, P., Yli-Juuti, T., Carslaw, K. S., Kulmala, M., Curtius, J., Baltensperger, U., Worsnop, D. R., Vehkamäki, H. and Kirkby, J.: Molecular understanding of sulphuric acid-amine particle nucleation in the atmosphere, Nature, 502(7471), 359–363,
- doi:10.1038/nature12663, 2013.
 Baker, A. K., Beyersdorf, A. J., Doezema, L. A., Katzenstein, A., Meinardi, S., Simpson, I. J., Blake, D. R. and Sherwood Rowland, F.: Measurements of nonmethane hydrocarbons in 28 United States cities, Atmos. Environ.,
- Boynard, A., Borbon, A., Leonardis, T., Barletta, B., Meinardi, S., Blake, D. R. and Locoge, N.: Spatial and seasonal variability
- of measured anthropogenic non-methane hydrocarbons in urban atmospheres: Implication on emission ratios, Atmos. Environ., doi:10.1016/j.atmosenv.2013.09.039, 2014.
 - Brean, J., Beddows, D. C. S., Shi, Z., Temime-Roussel, B., Marchand, N., Querol, X., Alastuey, A., Minguillón, M. C. and Harrison, R. M.: Molecular insights into new particle formation in Barcelona, Spain, Atmos. Chem. Phys., 20(16), 10029–10045, doi:10.5194/acp-20-10029-2020, 2020.
- 440 Breitenlechner, M., Fischer, L., Hainer, M., Heinritzi, M., Curtius, J. and Hansel, A.: PTR3: An instrument for studying the lifecycle of reactive organic carbon in the atmosphere, Anal. Chem., 89(11), 5824–5831, doi:10.1021/acs.analchem.6b05110, 2017.
 - Bruns, E. A., El Haddad, I., Slowik, J. G., Kilic, D., Klein, F., Baltensperger, U. and Prévôt, A. S. H.: Identification of significant precursor gases of secondary organic aerosols from residential wood combustion, Sci. Rep., doi:10.1038/srep27881,
- 445 2016.
 - Cai, R., Yang, D., Fu, Y., Wang, X., Li, X., Ma, Y., Hao, J., Zheng, J. and Jiang, J.: Aerosol surface area concentration: a governing factor in new particle formation in Beijing, Atmos. Chem. Phys., 17(20), 12327–12340, doi:10.5194/acp-17-12327-2017, 2017.
 - Cai, R., Yan, C., Yang, D., Yin, R., Lu, Y., Deng, C., Fu, Y., Ruan, J., Li, X., Kontkanen, J., Zhang, Q., Kangasluoma, J., Ma,
- 450 Y., Hao, J., Worsnop, D. R., Bianchi, F., Paasonen, P., Kerminen, V., Liu, Y., Wang, L., Zheng, J., Kulmala, M. and Jiang, J.: Sulfuric acid–amine nucleation in urban Beijing, Atmos. Chem. Phys., 21(4), 2457–2468, doi:10.5194/acp-21-2457-2021,

- 2021.
- Carnerero, C., Pérez, N., Reche, C., Ealo, M., Titos, G., Lee, H. K., Eun, H. R., Park, Y. H., Dada, L., Paasonen, P., Kerminen, V. M., Mantilla, E., Escudero, M., Gómez-Moreno, F. J., Alonso-Blanco, E., Coz, E., Saiz-Lopez, A., Temime-Roussel, B.,
- Marchand, N., Beddows, D. C. S., Harrison, R. M., Petäjä, T., Kulmala, M., Ahn, K. H., Alastuey, A. and Querol, X.: Vertical and horizontal distribution of regional new particle formation events in Madrid, Atmos. Chem. Phys., 18(22), 16601–16618, doi:10.5194/acp-18-16601-2018, 2018.
 - Chan, A. W. H., Kautzman, K. E., Chhabra, P. S., Surratt, J. D., Chan, M. N., Crounse, J. D., Kürten, A., Wennberg, P. O., Flagan, R. C. and Seinfeld, J. H.: Secondary organic aerosol formation from photooxidation of naphthalene and
- alkylnaphthalenes: implications for oxidation of intermediate volatility organic compounds (IVOCs), Atmos. Chem. Phys., 9(9), 3049–3060, doi:10.5194/acp-9-3049-2009, 2009.
 - Chan, T. W. and Mozurkewich, M.: Measurement of the coagulation rate constant for sulfuric acid particles as a function of particle size using tandem differential mobility analysis, J. Aerosol Sci., 32(3), 321–339, doi:10.1016/S0021-8502(00)00081-1, 2001.
- Dada, L., Lehtipalo, K., Kontkanen, J., Nieminen, T., Baalbaki, R., Ahonen, L., Duplissy, J., Yan, C., Chu, B., Petäjä, T., Lehtinen, K., Kerminen, V. M., Kulmala, M. and Kangasluoma, J.: Formation and growth of sub-3-nm aerosol particles in experimental chambers, Nat. Protoc., 15(3), 1013–1040, doi:10.1038/s41596-019-0274-z, 2020.
 - Daellenbach, K. R., Uzu, G., Jiang, J., Cassagnes, L.-E., Leni, Z., Vlachou, A., Stefenelli, G., Canonaco, F., Weber, S., Segers, A., Kuenen, J. J. P., Schaap, M., Favez, O., Albinet, A., Aksoyoglu, S., Dommen, J., Baltensperger, U., Geiser, M., El Haddad,
- 470 I., Jaffrezo, J.-L. and Prévôt, A. S. H.: Sources of particulate-matter air pollution and its oxidative potential in Europe, Nature, 587(7834), 414–419, doi:10.1038/s41586-020-2902-8, 2020.
 - Dai, L., Wang, H., Zhou, L., An, J., Tang, L., Lu, C., Yan, W., Liu, R., Kong, S., Chen, M., Lee, S. and Yu, H.: Regional and local new particle formation events observed in the Yangtze River Delta region, China, J. Geophys. Res., 122(4), 2389–2402, doi:10.1002/2016JD026030, 2017.
- Dunne, E. M., Gordon, H., Kurten, A., Almeida, J., Duplissy, J., Williamson, C., Ortega, I. K., Pringle, K. J., Adamov, A.,
 Baltensperger, U., Barmet, P., Benduhn, F., Bianchi, F., Breitenlechner, M., Clarke, A., Curtius, J., Dommen, J., Donahue, N.
 M., Ehrhart, S., Flagan, R. C., Franchin, A., Guida, R., Hakala, J., Hansel, A., Heinritzi, M., Jokinen, T., Kangasluoma, J.,
 Kirkby, J., Kulmala, M., Kupc, A., Lawler, M. J., Lehtipalo, K., Makhmutov, V., Mann, G., Mathot, S., Merikanto, J.,
 Miettinen, P., Nenes, A., Onnela, A., Rap, A., Reddington, C. L. S., Riccobono, F., Richards, N. A. D., Rissanen, M. P., Rondo,
- 480 L., Sarnela, N., Schobesberger, S., Sengupta, K., Simon, M., Sipila, M., Smith, J. N., Stozkhov, Y., Tome, A., Trostl, J., Wagner, P. E., Wimmer, D., Winkler, P. M., Worsnop, D. R. and Carslaw, K. S.: Global atmospheric particle formation from CERN CLOUD measurements, Science, 354(6316), 1119–1124, doi:10.1126/science.aaf2649, 2016.
 - Elser, M., El-Haddad, I., Maasikmets, M., Bozzetti, C., Wolf, R., Ciarelli, G., Slowik, J. G., Richter, R., Teinemaa, E., Hüglin, C., Baltensperger, U. and Prévôt, A. S. H.: High contributions of vehicular emissions to ammonia in three European cities
- derived from mobile measurements, Atmos. Environ., 175, 210–220, doi:10.1016/j.atmosenv.2017.11.030, 2018.

- Guo, H., Weber, R. J. and Nenes, A.: High levels of ammonia do not raise fine particle pH sufficiently to yield nitrogen oxide-dominated sulfate production, Sci. Rep., 7(1), 12109, doi:10.1038/s41598-017-11704-0, 2017.
- Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M., Zeng, L., Molina, M. J. and Zhang, R.: Elucidating severe urban haze formation in China, Proc. Natl. Acad. Sci. U. S. A., 111(49), 17373–17378, doi:10.1073/pnas.1419604111.2014.
 - Guo, S., Hu, M., Peng, J., Wu, Z., Zamora, M. L., Shang, D., Du, Z., Zheng, J., Fang, X., Tang, R., Wu, Y., Zeng, L., Shuai, S., Zhang, W., Wang, Y., Ji, Y., Li, Y., Zhang, A. L., Wang, W., Zhang, F., Zhao, J., Gong, X., Wang, C., Molina, M. J. and Zhang, R.: Remarkable nucleation and growth of ultrafine particles from vehicular exhaust, Proc. Natl. Acad. Sci. U. S. A., 117(7), 3427–3432, doi:10.1073/pnas.1916366117, 2020.
- Jathar, S. H., Gordona, T. D., Hennigan, C. J., Pye, H. O. T., Pouliot, G., Adams, P. J., Donahue, N. M. and Robinson, A. L.: Unspeciated organic emissions from combustion sources and their influence on the secondary organic aerosol budget in the United States, Proc. Natl. Acad. Sci. U. S. A., doi:10.1073/pnas.1323740111, 2014.
 Jayaratne, R., Pushpawela, B., He, C., Li, H., Gao, J., Chai, F. and Morawska, L.: Observations of particles at their formation sizes in Beijing, China, Atmos. Chem. Phys., 17, 8825–8835, doi:10.5194/acp-17-8825-2017, 2017.
- Jokinen, T., Sipilä, M., Junninen, H., Ehn, M., Lönn, G., Hakala, J., Petäjä, T., Mauldin, R. L., Kulmala, M. and Worsnop, D. R.: Atmospheric sulphuric acid and neutral cluster measurements using CI-APi-TOF, Atmos. Chem. Phys., 12(9), 4117–4125, doi:10.5194/acp-12-4117-2012, 2012.
 - Junninen, H., Ehn, M., Petäjä, Luosujärvi, L., Kotiaho, T., Kostiainen, R., Rohner, U., Gonin, M., Fuhrer, K., Kulmala, M. and Worsnop, D. R.: A high-resolution mass spectrometer to measure atmospheric ion composition, Atmos. Meas. Tech., 3(4),
- Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné, S., Ickes, L., Kürten, A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G., Wimmer, D., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud, W., Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R., Makhmutov, V., Mathot, S.,

1039–1053, doi:10.5194/amt-3-1039-2010, 2010.

- Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A., Pereira, P., Petäjä, T., Schnitzhofer, R., Seinfeld, J. H., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A., Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U. and Kulmala, M.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation, Nature, 476(7361), 429–435, doi:10.1038/nature10343, 2011.
- Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M., Yan, C., Almeida, J., Tröstl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A., Bernhammer, A.-K., Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X., Craven, J., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim, J., Krapf, M., Kürten, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Peräkylä, O., Piel, F., Petäjä, T., Praplan, A. P., Pringle, K., Rap, A., Richards, N. A.

- D., Riipinen, I., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Scott, C. E., Seinfeld, J. H., Sipilä, M., Steiner, G., Stozhkov, Y., Stratmann, F., Tomé, A., Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R., Baltensperger, U., Kulmala, M., Carslaw, K. S. and Curtius, J.: Ion-induced nucleation of pure biogenic particles, Nature, 533(7604), 521–526, doi:10.1038/nature17953, 2016.
- Kontkanen, J., Lehtipalo, K., Ahonen, L., Kangasluoma, J., Manninen, H. E., Hakala, J., Rose, C., Sellegri, K., Xiao, S., Wang, L., Qi, X., Nie, W., Ding, A., Yu, H., Lee, S., Kerminen, V.-M., Petäjä, T. and Kulmala, M.: Measurements of sub-3 nm particles using a particle size magnifier in different environments: from clean mountain top to polluted megacities, Atmos. Chem. Phys., 17, 2163–2187, doi:10.5194/acp-17-2163-2017, 2017.
- Kuang, C., McMurry, P. H., McCormick, A. V. and Eisele, F. L.: Dependence of nucleation rates on sulfuric acid vapor concentration in diverse atmospheric locations, J. Geophys. Res., 113(D10), D10209, doi:10.1029/2007JD009253, 2008.
- Kuang, C., Riipinen, I., Sihto, S.-L., Kulmala, M., Mccormick, A. V and Mcmurry, P. H.: An improved criterion for new particle formation in diverse atmospheric environments, Atmos. Chem. Phys., 10, 8469–8480, doi:10.5194/acp-10-8469-2010, 2010.
- Kulmala, M., Kerminen, V. M., Petäjä, T., Ding, A. J. and Wang, L.: Atmospheric gas-to-particle conversion: Why NPF events are observed in megacities?, Faraday Discuss., 200, 271–288, doi:10.1039/c6fd00257a, 2017.
 - Kürten, A.: New particle formation from sulfuric acid and ammonia: Nucleation and growth model based on thermodynamics derived from CLOUD measurements for a wide range of conditions, Atmos. Chem. Phys., 19(7), 5033–5050, doi:10.5194/acp-19-5033-2019, 2019.
 - Kürten, A., Bianchi, F., Almeida, J., Kupiainen-Määttä, O., Dunne, E. M., Duplissy, J., Williamson, C., Barmet, P.,
- Breitenlechner, M., Dommen, J., Donahue, N. M., Flagan, R. C., Franchin, A., Gordon, H., Hakala, J., Hansel, A., Heinritzi, M., Ickes, L., Jokinen, T., Kangasluoma, J., Kim, J., Kirkby, J., Kupc, A., Lehtipalo, K., Leiminger, M., Makhmutov, V., Onnela, A., Ortega, I. K., Petäjä, T., Praplan, A. P., Riccobono, F., Rissanen, M. P., Rondo, L., Schnitzhofer, R., Schobesberger, S., Smith, J. N., Steiner, G., Stozhkov, Y., Tomé, A., Tröstl, J., Tsagkogeorgas, G., Wagner, P. E., Wimmer, D., Ye, P., Baltensperger, U., Carslaw, K., Kulmala, M. and Curtius, J.: Experimental particle formation rates spanning
- tropospheric sulfuric acid and ammonia abundances, ion production rates, and temperatures, J. Geophys. Res., 121(20), 12,377-12,400, doi:10.1002/2015JD023908, 2016.
 - Kürten, A., Li, C., Bianchi, F., Curtius, J., Dias, A., Donahue, N. M., Duplissy, J., Flagan, R. C., Hakala, J., Jokinen, T., Kirkby, J., Kulmala, M., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Onnela, A., Rissanen, M. P., Simon, M., Sipilä, M., Stozhkov, Y., Tröstl, J., Ye, P. and McMurry, P. H.: New particle formation in the sulfuric acid-dimethylamine-water system:
- Reevaluation of CLOUD chamber measurements and comparison to an aerosol nucleation and growth model, Atmos. Chem. Phys., 18(2), 845–863, doi:10.5194/acp-18-845-2018, 2018.
 - Lehtipalo, K., Leppä, J., Kontkanen, J., Kangasluoma, J., Franchini, A., Wimmer, D., Schobesberger, S., Junninen, H., Petäjä, T., Sipilä, M., Mikkilä, J., Vanhanen, J., Worsnop, D. R. and Kulmala, M.: Methods for determining particle size distribution

- and growth rates between 1 and 3 nm using the Particle Size Magnifier, Boreal Environ. Res., 19(B), 15-236, 2014.
- Lehtipalo, K., Rondo, L., Kontkanen, J., Schobesberger, S., Jokinen, T., Sarnela, N., Kürten, A., Ehrhart, S., Franchin, A., Nieminen, T., Riccobono, F., Sipilä, M., Yli-Juuti, T., Duplissy, J., Adamov, A., Ahlm, L., Almeida, J., Amorim, A., Bianchi, F., Breitenlechner, M., Dommen, J., Downard, A. J., Dunne, E. M., Flagan, R. C., Guida, R., Hakala, J., Hansel, A., Jud, W., Kangasluoma, J., Kerminen, V. M., Keskinen, H., Kim, J., Kirkby, J., Kupc, A., Kupiainen-Määttä, O., Laaksonen, A., Lawler, M. J., Leiminger, M., Mathot, S., Olenius, T., Ortega, I. K., Onnela, A., Petäjä, T., Praplan, A., Rissanen, M. P., Ruuskanen,
- T., Santos, F. D., Schallhart, S., Schnitzhofer, R., Simon, M., Smith, J. N., Tröstl, J., Tsagkogeorgas, G., Tomé, A., Vaattovaara, P., Vehkamäki, H., Vrtala, A. E., Wagner, P. E., Williamson, C., Wimmer, D., Winkler, P. M., Virtanen, A., Donahue, N. M., Carslaw, K. S., Baltensperger, U., Riipinen, I., Curtius, J., Worsnop, D. R. and Kulmala, M.: The effect of acid–base clustering and ions on the growth of atmospheric nano-particles, Nat. Commun., 7(1), 11594, doi:10.1038/ncomms11594, 2016.
- Lehtipalo, K., Yan, C., Dada, L., Bianchi, F., Xiao, M., Wagner, R., Stolzenburg, D., Ahonen, L. R., Amorim, A., Baccarini, A., Bauer, P. S., Baumgartner, B., Bergen, A., Bernhammer, A. K., Breitenlechner, M., Brilke, S., Buchholz, A., Mazon, S. B., Chen, D., Chen, X., Dias, A., Dommen, J., Draper, D. C., Duplissy, J., Ehn, M., Finkenzeller, H., Fischer, L., Frege, C., Fuchs, C., Garmash, O., Gordon, H., Hakala, J., He, X., Heikkinen, L., Heinritzi, M., Helm, J. C., Hofbauer, V., Hoyle, C. R., Jokinen, T., Kangasluoma, J., Kerminen, V. M., Kim, C., Kirkby, J., Kontkanen, J., Kürten, A., Lawler, M. J., Mai, H., Mathot,
- 570 S., Mauldin, R. L., Molteni, U., Nichman, L., Nie, W., Nieminen, T., Ojdanic, A., Onnela, A., Passananti, M., Petäjä, T., Piel, F., Pospisilova, V., Quéléver, L. L. J., Rissanen, M. P., Rose, C., Sarnela, N., Schallhart, S., Schuchmann, S., Sengupta, K., Simon, M., Sipilä, M., Tauber, C., Tomé, A., Tröstl, J., Väisänen, O., Vogel, A. L., Volkamer, R., Wagner, A. C., Wang, M., Weitz, L., Wimmer, D., Ye, P., Ylisirniö, A., Zha, Q., Carslaw, K. S., Curtius, J., Donahue, N. M., Flagan, R. C., Hansel, A., Riipinen, I., Virtanen, A., Winkler, P. M., Baltensperger, U., Kulmala, M. and Worsnop, D. R.: Multicomponent new particle
- formation from sulfuric acid, ammonia, and biogenic vapors, Sci. Adv., 4(12), eaau5363, doi:10.1126/sciadv.aau5363, 2018. Li, L., Tang, P., Nakao, S. and Cocker III, D. R.: Impact of molecular structure on secondary organic aerosol formation from aromatic hydrocarbon photooxidation under low-NOx conditions, Atmos. Chem. Phys., 16(17), 10793–10808, doi:10.5194/acp-16-10793-2016, 2016.
- Li, M., Zhang, Q., Zheng, B., Tong, D., Lei, Y., Liu, F., Hong, C., Kang, S., Yan, L., Zhang, Y., Bo, Y., Su, H., Cheng, Y. and He, K.: Persistent growth of anthropogenic non-methane volatile organic compound (NMVOC) emissions in China during 1990-2017: Drivers, speciation and ozone formation potential, Atmos. Chem. Phys., doi:10.5194/acp-19-8897-2019, 2019. Mirme, S. and Mirme, A.: The mathematical principles and design of the NAIS A spectrometer for the measurement of cluster ion and nanometer aerosol size distributions, Atmos. Meas. Tech., 6(4), 1061–1071, doi:10.5194/amt-6-1061-2013, 2013.
- Molteni, U., Bianchi, F., Klein, F., Haddad, I. El, Frege, C., Rossi, M. J., Dommen, J. and Baltensperger, U.: Formation of highly oxygenated organic molecules from aromatic compounds, Atmos. Chem. Phys., 18, 1909–1921, doi:10.5194/acp-18-1909-2018, 2018.

- Myllys, N., Kubečka, J., Besel, V., Alfaouri, D., Olenius, T., Smith, J. N. and Passananti, M.: Role of base strength, cluster structure and charge in sulfuric-acid-driven particle formation, Atmos. Chem. Phys., 19(15), 9753–9768, doi:10.5194/acp-19-9753-2019, 2019.
 - Nieminen, T., Lehtinen, K. E. J. and Kulmala, M.: Sub-10 nm particle growth by vapor condensation effects of vapor molecule size and particle thermal speed, Atmos. Chem. Phys., 10, 9773–9779, doi:10.5194/acp-10-9773-2010, 2010.
 - Paasonen, P., Nieminen, T., Asmi, E., Manninen, H. E., Petäjä, T., Plass-Dülmer, C., Flentje, H., Birmili, W., Wiedensohler, A., Hõrrak, U., Metzger, A., Hamed, A., Laaksonen, A., Facchini, M. C., Kerminen, V. M. and Kulmala, M.: On the roles of
- sulphuric acid and low-volatility organic vapours in the initial steps of atmospheric new particle formation, Atmos. Chem. Phys., 10(22), 11223–11242, doi:10.5194/acp-10-11223-2010, 2010.
 - Pfeifer, J., Simon, M., Heinritzi, M., Piel, F., Weitz, L., Wang, D., Granzin, M., Müller, T., Bräkling, S., Kirkby, J., Curtius, J. and Kürten, A.: Measurement of ammonia, amines and iodine compounds using protonated water cluster chemical ionization mass spectrometry, Atmos. Meas. Tech., 13(5), 2501–2522, doi:10.5194/amt-13-2501-2020, 2020.
- Riccobono, F., Schobesberger, S., Scott, C. E., Dommen, J., Ortega, I. K., Rondo, L., Almeida, J., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Downard, A., Dunne, E. M., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Hansel, A., Junninen, H., Kajos, M., Keskinen, H., Kupc, A., Kürten, A., Kvashin, A. N., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Petäjä, T., Praplan, A. P., Santos, F. D., Schallhart, S., Seinfeld, J. H., Sipilä, M., Spracklen, D. V, Stozhkov, Y., Stratmann, F., Tomé, A., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y., Vrtala, A., Wagner,
- P. E., Weingartner, E., Wex, H., Wimmer, D., Carslaw, K. S., Curtius, J., Donahue, N. M., Kirkby, J., Kulmala, M., Worsnop, D. R. and Baltensperger, U.: Oxidation products of biogenic emissions contribute to nucleation of atmospheric particles, Science, 344(6185), 717–721, doi:10.1126/science.1243527, 2014.
 - Seinfeld, J. H. and Pandis, S. N.: Atmospheric Chemistry and Physics: From Air Pollution to Climate Change, 3rd Editio., John Wiley & Sons, Ltd., 2016.
- Simon, M., Heinritzi, M., Herzog, S., Leiminger, M., Bianchi, F., Praplan, A., Dommen, J., Curtius, J. and Kurten, A.: Detection of dimethylamine in the low pptv range using nitrate chemical ionization atmospheric pressure interface time-of-flight (CI-APi-TOF) mass spectrometry, Atmos. Meas. Tech., 9(5), 2135–2145, doi:10.5194/amt-9-2135-2016, 2016.
 Stolzenburg, D., Steiner, G. and Winkler, P. M.: A DMA-Train for precision measurement of sub-10nm aerosol dynamics, Atmos. Meas. Tech., 10(4), 1639–1651, doi:10.5194/amt-10-1639-2017, 2017.
- Stolzenburg, D., Stolzenburg, D., Simon, M., Ranjithkumar, A., Kürten, A., Lehtipalo, K., Lehtipalo, K., Gordon, H., Ehrhart, S., Finkenzeller, H., Pichelstorfer, L., Nieminen, T., He, X. C., Brilke, S., Xiao, M., Amorim, A., Baalbaki, R., Baccarini, A., Beck, L., Bräkling, S., Murillo, L. C., Chen, D., Chu, B., Dada, L., Dias, A., Dommen, J., Duplissy, J., El Haddad, I., Fischer, L., Carracedo, L. G., Heinritzi, M., Kim, C., Kim, C., Koenig, T. K., Kong, W., Lamkaddam, H., Lee, C. P., Leiminger, M., Leiminger, M., Li, Z., Makhmutov, V., Manninen, H. E., Marie, G., Marten, R., Müller, T., Nie, W., Partoll, E., Petäjä, T.,
- Pfeifer, J., Philippov, M., Rissanen, M. P., Rissanen, M. P., Rörup, B., Schobesberger, S., Schuchmann, S., Shen, J., Sipilä,
 M., Steiner, G., Stozhkov, Y., Tauber, C., Tham, Y. J., Tomé, A., Vazquez-Pufleau, M., Wagner, A. C., Wagner, A. C., Wang,

- M., Wang, Y., Weber, S. K., Wimmer, D., Wimmer, D., Wlasits, P. J., Wu, Y., Ye, Q., Zauner-Wieczorek, M., Baltensperger, U., Carslaw, K. S., Curtius, J., Donahue, N. M., Flagan, R. C., Hansel, A., Hansel, A., Kulmala, M., Lelieveld, J., Volkamer, R., Kirkby, J., Kirkby, J. and Winkler, P. M.: Enhanced growth rate of atmospheric particles from sulfuric acid, Atmos. Chem.
- 625 Phys., 20(12), 7359–7372, doi:10.5194/acp-20-7359-2020, 2020.
 - Tröstl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege, C., Bianchi, F., Wagner, R., Simon, M., Lehtipalo, K., Williamson, C., Craven, J. S., Duplissy, J., Adamov, A., Almeida, J., Bernhammer, A.-K., Breitenlechner, M., Brilke, S., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Gysel, M., Hansel, A., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Keskinen, H., Kim, J., Krapf, M., Kürten, A., Laaksonen, A., Lawler, M.,
- 630 Leiminger, M., Mathot, S., Möhler, O., Nieminen, T., Onnela, A., Petäjä, T., Piel, F. M., Miettinen, P., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Sengupta, K., Sipilä, M., Smith, J. N., Steiner, G., Tomè, A., Virtanen, A., Wagner, A. C., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Carslaw, K. S., Curtius, J., Dommen, J., Kirkby, J., Kulmala, M., Riipinen, I., Worsnop, D. R., Donahue, N. M. and Baltensperger, U.: The role of low-volatility organic compounds in initial particle growth in the atmosphere, Nature, 533(7604), 527–531, doi:10.1038/nature18271, 2016.
- Vanhanen, J., Mikkilä, J., Lehtipalo, K., Sipilä, M., Manninen, H. E., Siivola, E., Petäjä, T. and Kulmala, M.: Particle size magnifier for nano-CN detection, Aerosol Sci. Technol., 45(4), 533–542, doi:10.1080/02786826.2010.547889, 2011.
 Volkamer, R., Volkamer, R., Volkamer, R., Klotz, B., Klotz, B., Barnes, I., Imamura, T., Wirtz, K., Washida, N., Becker, K. H. and Platt, U.: OH-initiated oxidation of benzene: Part I. Phenol formation under atmospheric conditions, Phys. Chem. Chem. Phys., 4(9), 1598–1610, doi:10.1039/b108747a, 2002.
- Wang, M., Chen, D., Xiao, M., Ye, Q., Stolzenburg, D., Hofbauer, V., Ye, P., Vogel, A. L., Mauldin, R. L., Amorim, A., Baccarini, A., Baumgartner, B., Brilke, S., Dada, L., Dias, A., Duplissy, J., Finkenzeller, H., Garmash, O., He, X.-C., Hoyle, C. R., Kim, C., Kvashnin, A., Lehtipalo, K., Fischer, L., Molteni, U., Petäjä, T., Pospisilova, V., Quéléver, L. L. J., Rissanen, M., Simon, M., Tauber, C., Tomé, A., Wagner, A. C., Weitz, L., Volkamer, R., Winkler, P. M., Kirkby, J., Worsnop, D. R., Kulmala, M., Baltensperger, U., Dommen, J., El-Haddad, I. and Donahue, N. M.: Photo-oxidation of aromatic hydrocarbons
- 645 produces low-volatility organic compounds, Environ. Sci. Technol., 54(13), 7911–7921, doi:10.1021/acs.est.0c02100, 2020.
 Wang, S., Wu, R., Berndt, T., Ehn, M. and Wang, L.: Formation of Highly Oxidized Radicals and Multifunctional Products from the Atmospheric Oxidation of Alkylbenzenes, Environ. Sci. Technol., 51(15), 8442–8449, doi:10.1021/acs.est.7b02374, 2017.
- Yao, L., Garmash, O., Bianchi, F., Zheng, J., Yan, C., Kontkanen, J., Junninen, H., Mazon, S. B., Ehn, M., Paasonen, P., Sipilä,
 M., Wang, M., Wang, X., Xiao, S., Chen, H., Lu, Y., Zhang, B., Wang, D., Fu, Q., Geng, F., Li, L., Wang, H., Qiao, L., Yang, X., Chen, J., Kerminen, V. M., Petäjä, T., Worsnop, D. R., Kulmala, M. and Wang, L.: Atmospheric new particle formation from sulfuric acid and amines in a Chinese megacity, Science, 361(6399), 278–281, doi:10.1126/science.aao4839, 2018.
 - Yu, H., Zhou, L., Dai, L., Shen, W., Dai, W., Zheng, J., Ma, Y. and Chen, M.: Nucleation and growth of sub-3 nm particles in the polluted urban atmosphere of a megacity in China, Atmos. Chem. Phys., 16, 2641–2657, doi:10.5194/acp-16-2641-2016,
- 655 2016.

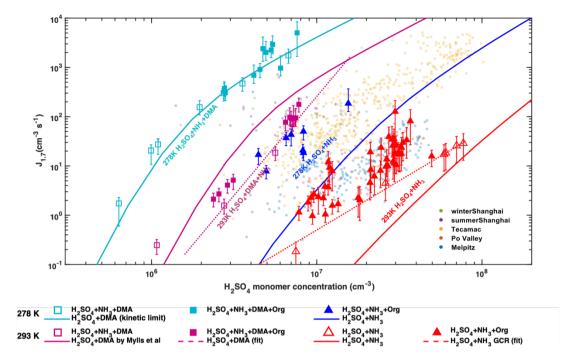
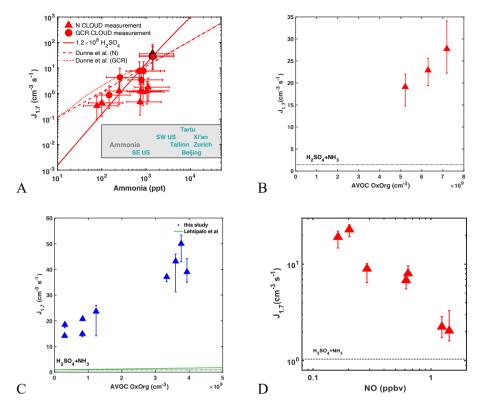


Figure 1: Dependencies of atmospheric particle formation rates (J_{1.7}) on sulfuric acid. J_{1.7} versus H₂SO₄ measured in CLOUD at 278 K (blue and cyan) or 293 K (red and magenta) with 1 to 2 ppbv NH₃. Experiments without DMA injection are shown by triangles and those with 4 pptv DMA by squares. Filled and open symbols indicate presence or absence of anthropogenic organics, respectively, where the variability of the filled symbols is mostly due to different concentrations of oxidised organics (OxOrg) from NAPH, TMB and TOL oxidation, and NO_x. Over the range of 10⁶ to 10⁸ cm⁻³ sulfuric acid, J_{1.7} increases from 1 to 10⁴ cm⁻³s⁻¹. The error bars indicate the measurement uncertainty of the nucleation rates. Atmospheric observations in the polluted boundary layer are indicated by small coloured circles (Kuang et al., 2008; Paasonen et al., 2010; Yao et al., 2018). Solid lines show the predicted nucleation rates of sulfuric acid and 2 ppbv ammonia at 278 K (blue) and 293 K (red) from the kinetic model. The nucleation rate of H₂SO₄-NH₃ under GCR conditions is given by the red dashed line. The nucleation rate of H₂SO₄ at the kinetic limit is indicated by the solid cyan curve, which matches our H₂SO₄+DMA experiments The nucleation rate of H₂SO₄ at the kinetic limit is indicated by the solid cyan curve. The solid magenta line is the predicted nucleation rates of sulfuric acid and 4 pptv DMA at 293K and the magenta dashed line shows a fit of J_{1.7} = k [H₂SO₄]⁴ to the data with 4 pptv DMA at 293K.



680

685

Figure 2: Dependencies of atmospheric particle formation rates (J_{17}) on ammonia, organics and NO. (A) J_{17} versus NH₃ at fixed H₂SO₄ ((1.1-1.6) × 10⁸ cm⁻³) and temperature (292-295 K). Red triangles are previous CLOUD measurements (Dunne et al., 2016) under neutral (triangle) or GCR (circle) conditions while the symbols with black outlines are new from this study. Error bars describe run-to-run repeatability of J in CLOUD, which is ca. 30%. The solid red line presents J_{L7} for 1.2×10^8 cm⁻³ H₂SO₄ at 293 K from our kinetic model simulation and fits well to the additional data at the high NH₃ mixing ratio. The dashed (N, neutral) or dotted (GCR) red line shows the parameterisation used by Dunne et al. (2016). Also shown are the mean concentrations of ammonia at various locations (Elser et al., 2018; Guo et al., 2017) as text of locations. (B) Experiments with $2.8\pm0.3\times10^7$ cm⁻³ H₂SO₄, 849 ± 20 pptv NH₃ and 0.19 ± 0.03 ppbv NO at 293 K with NAPH:TMB:TOL of 1:2:10. Black dotted line in the plot displays J_{L7} for 1.9×10^7 H₂SO₄ nucleating with 1000 pptv NH₃. (C) Experiments with $8\pm1.5\times10^6$ cm⁻³ H₂SO₄, 1350±350 pptv NH₃ and 0.51±0.3 ppbv NO at 278 K, with NAPH:TMB:TOL of 1:5:30. Black dotted line in the plot displays J_{L7} for 8×10^6 cm⁻³ H₂SO₄ nucleating with 1600 pptv NH₃. Green solid line shows J_{L7} of 8×10^6 cm⁻³ H₂SO₄ nucleating with 1600 pptv NH₃, with BVOC OxOrg and with 0.2 ppbv NO based on parameterizations in Lehtipalo et al. (2018). Oxidation of aromatic compounds by OH forms highly oxygenated molecules (HOMs) (Molteni et al., 2018; Wang et al., 2017) through autoxidation, leading to significant NPF, at rates even higher than observed for monoterpene oxidation. This is consistent with the finding that HOMs from aromatic compounds are less volatile than HOMs from monoterpene precursors at the same number of carbon atoms in the backbone (Wang et al., 2020). (D) J_{L7} versus NO at 293K, with fixed conditions for other vapours $(2.3\pm0.4) \times 10^7$ cm⁻³ H₂SO₄, 800 to 1000 pptv ammonia, and an AVOC mixture yielding an OxOrg concentration of (5.2±0.7) ×109 cm⁻³ (among which (0.9±0.2) ×109 cm⁻³ produced by NAPH oxidation, $(2.1\pm0.3)\times10^9$ cm⁻³ by TMB oxidation and $(2.2\pm0.3)\times10^9$ cm⁻³ by TOL oxidation). The suppression of $J_{1.7}$ with increasing NO is due to reduced production of extremely low volatility OxOrg. The error bars describe the variation of the measurements only, and do not include a 30% run-to-run variability. The horizontal dotted line indicates J_{L7} for 2.3×10^7 cm⁻³ H_2SO_4 with 1000 pptv NH₃ at 293 K under GCR conditions. $J_{1,7}$ decreases with increasing NO since NO suppresses both autoxidation and the formation of low-volatility dimers. This efficient suppression of aromatic RO₂ isomerization to form autoxidation HOM products is in contrast to aromatic RO₂ isomerization rates to form bicycloalkyl radicals, which dominate the early stage of aromatic oxidation and outcompete NO reactions even at much higher NO concentrations (10s to 100s of ppbv) (Volkamer et al., 2002).

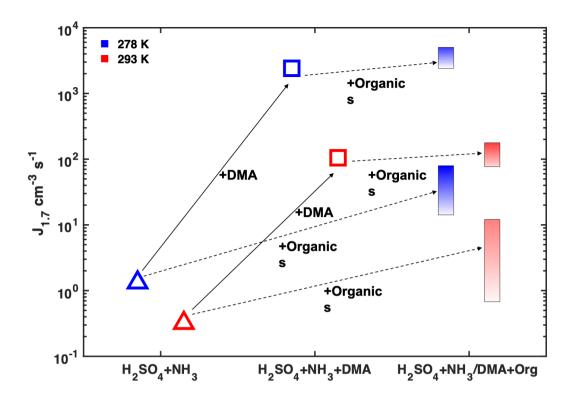
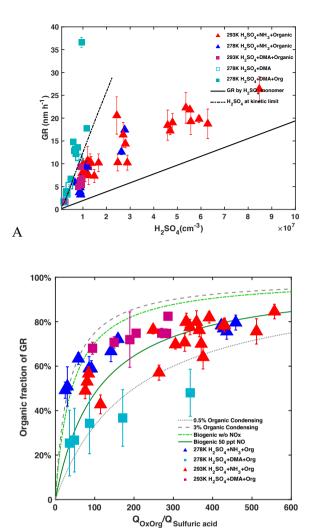


Figure 3. Figure 3: Simplified scheme of NPF rates in polluted environments showing the effect of adding different vapours on H₂SO₄-NH₃ nucleation at two different temperatures. All points have similar H₂SO₄ (8±2 × 10⁷ cm⁻³) and NH₃ (1-2 ppbv). The leftmost points are measured with only SO₂ and NH₃ added to the chamber, and each step to the right represents the addition of one more component to the system. Solid arrows describe the addition of ca. 4 pptv of DMA, dashed arrows describe the addition of aromatic hydrocarbons. Increasing color intensity of bars indicates increasing production rate of highly oxidised organics.



В 700 Figure 4: Influence of sulfuric acid, OxOrg and DMA on growth rate (GR): (A) GRs are measured between 1.8 nm and 3.2 nm, at 278 K (blue and cyan) or 293 K (red and magenta) with 1 to 2 ppbv NH₃ and plotted versus sulfuric acid monomer concentration. GRs without DMA injection are shown by triangles and GRs with 4 pptv DMA by squares. Filled and open symbols indicate presence or absence of organics, respectively, where the variability of the filled symbols is mostly driven by the concentrations of oxidised organics (OxOrg) and NO_x. The solid line shows GR for the condensation of sulfuric acid monomer only while the dashed line displays sulfuric acid nucleating at the kinetic limit, when cluster coagulation dominates the growth. (B) Contribution of oxidised organics (OxOrg) to sub-3 nm GR (determined by subtracting GR by sulfuric acid) as a function of the ratio of the production rates of oxidised organics (Q_{OxOrg}) to sulfuric acid ($Q_{sulfuric}$) acid). This ratio is roughly proportional to the ratio of condensable organics to sulfuric acid in the CLOUD chamber. The grey lines encompass most data and indicate that 0.5% (dotted) to 3% (dashed) of the total oxidised organics contribute to the growth of sub-3 nm particles. With the exception of sulfuric acid + DMA at 278 K, organics are the dominant contributor to particle growth in the sub-3 nm size range. At larger 710 particle sizes, the contribution of organics will increase further (Tröstl et al., 2016). Data are compared to the parameterisations of monoterpene oxidation by Lehtipalo et al. (2018) where the dark green dashdot line shows a strong decrease in the yield of condensable gases in the presence of only 50 pptv NO compared to the light green line for conditions without NO. Condensable OxOrg yields are generally higher from AVOCs than those obtained from biogenic vapours such as monoterpenes at elevated NO_X conditions (green lines), again consistent with the lower volatility of aromatic OxOrg formed by multi-generation oxidation in comparison with biogenic (monoterpene) 715 OxOrg with the same oxygen content (Wang et al., 2020a).

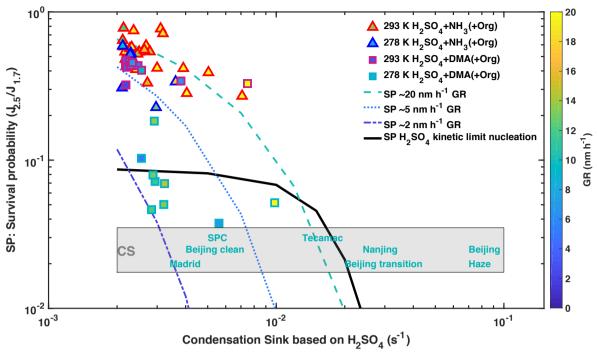


Figure 5: Influence of OxOrg and DMA on survival probability: Survival probability (SP) vs condensation sink (CS₆). The survival probability (SP) is calculated as the ratio between the formation rates at 2.5 nm and 1.7 nm. CLOUD data were recorded at 60 % RH and 278 K (blue and cyan contoured) or 293 K (red and magenta contoured) with 1 to 2 ppbv NH₃. Experiments without DMA injection are shown by triangles, and experiments with 4 pptv DMA by squares. The sub-3 nm growth rate is given by the colour of the symbols. The loss rates of newly formed particles are approximated by the condensation sink of sulfuric acid to particles larger than 6 nm (x-axis). Also shown are typical CS from observations in the polluted boundary layer at the following locations: Po Valley regional (San Pietro Capofiume, SPC) (Kontkanen et al., 2017), Madrid (Carnerero et al., 2018), Tecamac (Kuang et al., 2010), Nanjing (Yu et al., 2016) and Beijing (clean, transition and haze) (Cai et al., 2017)Condensation sinks observed in the atmosphere are indicated as text of the location (Cai et al., 2017; Carnerero et al., 2018; Kontkanen et al., 2016; Kuang et al., 2010; Yu et al., 2016). Dashed lines are calculated survival probabilities of 2.5 nm particles with growth rates of 2, 5, and 10 nm h⁻¹. The solid line is the survival probability of 2.5 nm particles if 8 x 10⁶ cm⁻³ sulfuric acid nucleates at the kinetic limit, which would e.g. at CS₆ = 0.002 s⁻¹ (representing wall loss) correspond to a nucleation rate J_{1.7} of 4000 cm⁻³ s⁻¹ (Fig. 1) and a growth rate of 11 nm h⁻¹ (Fig. 4A). The high production rate of nucleating clusters significantly reduces the survival probability, due to cluster-cluster collisions. However, at the same time, this increases the growth rate such that the survival probability is unaffected by the CS until it becomes comparable to the loss rate from cluster-cluster collisions.

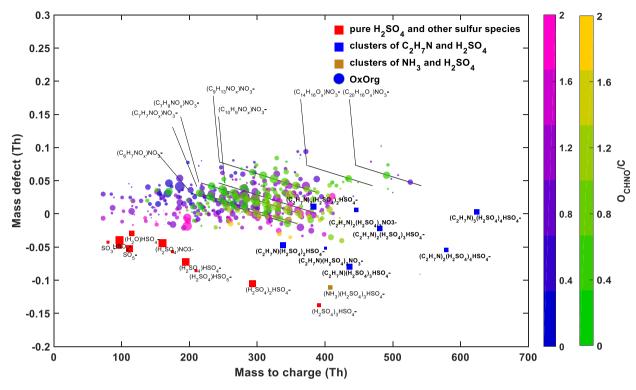


Figure 6. Mass defect plot of sulfuric acid nucleation with DMA in the presence of organics at 278 K in the CLOUD chamber. The conditions are 5.2 x 10⁶ cm⁻³ H₂SO₄ monomer, 4 pptv DMA, 1200 pptv NH₃, 680 pptv NO, 420 pptv, NAPH, 2200 pptv TMB and 9250 pptv TOL. Sulfate clusters are shown by red squares, sulfate-DMA clusters containing up to 5 sulfur molecules by blue squares, sulfate-NH₃ clusters by brown squares, and organics by circles coloured by O:C and by their chemical composition CHO or CHNO, respectively. The size of the symbols is proportional to their intensity in the mass spectrum.

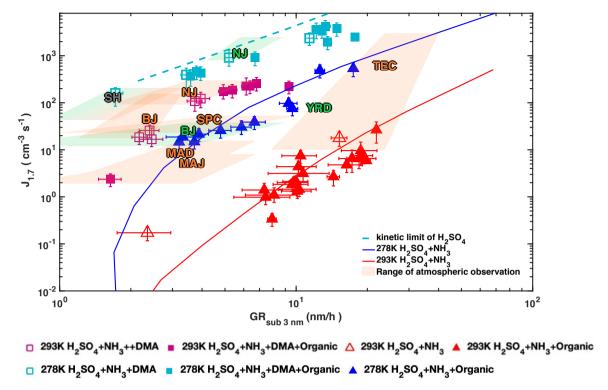


Figure 7: Particle formation rates (*J*_{1.7}) versus growth rates of sub-3 nm particles. CLOUD data were recorded at 278 K (blue and cyan symbols) or 293 K (red and magenta) with 1 to 2 ppbv NH₃. Experiments without DMA are shown by triangles and experiments with 4 pptv DMA by squares. Filled and open symbols indicate presence or absence of organics, respectively. Error bars denote uncertainties of *J*_{1.7} and GR values. Solid lines represent the simulated nucleation/growth rates with a kinetic model of sulfuric acid with 1 ppbv ammonia at 278 K (blue) and 293 K (red). H₂SO₄ nucleation at the kinetic limit is indicated by the cyan dashed line. Also shown are the mean values for sub-3 nm GR and sub-2 nm *J* from observations in the polluted boundary layer at the following locations: Shanghai (SH) (Yao et al., 2018), Beijing spring and winter(BJ) (Cai et al., 2017; Jayaratne et al., 2017), Madrid (urban: MAD and suburban: MAJ) (Carnerero et al., 2018), Po Valley regional (San Pietro Capofiume, SPC) (Kontkanen et al., 2017), Nanjing summer and winter (NJ) (Yu et al., 2016), Yangtze-River-Delta regional (YRD) (Dai et al., 2017) and Tecamac (TEC) (Kuang et al., 2008). The text colours indicate <288 K (cyan), >288 K (orange) or all year (grey). The shaded area colours indicate the range of (*J*, GR) values at these locations for temperatures <288 K (cyan) and >288 K (orange).