

Changes control version of **Influence of sea salt aerosols on the development of Mediterranean tropical-like cyclones**

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Abstract. Medicanes are mesoscale tropical-like cyclones that develop in the Mediterranean basin and represent a great hazard for the coastal population. The skill to accurately simulate them is of utmost importance to prevent economical and personal damages. Medicanes are fuelled by the latent heat released in the condensation process associated to convective activity, which is regulated by the presence and activation of cloud condensation nuclei, ~~originated mainly~~ mainly originated from sea salt aerosols (SSA) for marine environments. Henceforth, the purpose of this contribution is twofold: assessing the effects of an interactive calculation of SSA on the strengthening and persistence of medicanes; and providing insight on the casuistry and sensitivities around their simulation processes. To this end, a set of simulations has been conducted with a chemistry/meteorology coupled model considering prescribed aerosols (PA) and interactive aerosol concentrations (IA). The results indicate that IA produces longer-lasting and more intense medicanes. Further, the role of the initialization time and nudging strategies for medicane simulations has been explored. Overall, the results suggest that (1) the application of spectral nudging dampens the effects of IA; (2) the initialization time introduces a strong variability on the storm dynamics; and (3) wind-SSA feedback is crucial and should be considered when studying medicanes.

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1 Introduction

Mediterranean tropical-like cyclones, also known as medicanes (from **mediterranean hurricanes**), are mesoscale perturbations that exhibit tropical characteristics, such as an eye-like feature and warm core. These storms are characterized by high wind speeds and vertically aligned geopotential height perturbations along different pressure levels. Just like regular tropical cyclones, medicanes represent a hazard for the population of coastal areas. However, given the relatively small extent of the Mediterranean basin and the lower sea surface temperatures of the Mediterranean Sea, they do not reach the size and intensity of actual hurricanes. Still, they can produce heavy precipitation and intense wind gusts, reaching up to Category 1 in the

Saffir-Simpson scale (Fita et al., 2007; Miglietta and Rotunno, 2019). Thus, our ability to understand and ~~accurately simulate~~ simulate accurately medicanes with state-of-the-art meteorological modeling systems stands as a key factor to prevent their associated damages.

25 Tropical-like cyclones in general, and medicanes in particular, are usually considered to be a hybrid between tropical and extratropical cyclones. Although their triggering and development mechanisms differ from those of tropical cyclones (~~Tous and Romero, 2013; Cavicchia et al., 2014; Miglietta and Rotunno, 2019~~) (Tous and Romero, 2013; Cavicchia et al., 2014; Miglietta, such storms are ~~maintained in the same way~~ generally maintained as tropical cyclones, at least in their mature stage, in which moist and warm air is advected towards the low pressure storm centre at ~~surface and the~~ the surface and lower tropospheric levels. This produces and maintains a thermal disequilibrium in the vertical direction that enhances convection, consequently
30 producing condensation through adiabatic cooling. The condensation process releases latent heat that warms the cyclone core, which is the dominant mechanism that sustains the vortex structure (Lagouvardos et al., 1999).

Numerous studies have addressed the sensitivity of medicane simulations to different factors related to their intensification and track. Some authors have studied the effects of an increased sea surface temperature (SST) (Pytharoulis, 2018; Noyelle et al., 2019), while others focus on the role of air-sea interaction and surface heat fluxes (Tous et al., 2013; Akhtar et al., 2014;
35 Ricchi et al., 2017; Gaertner et al., 2018; Rizza et al., 2018; Bouin and Lebeau-pin Brossier, 2020) or the influence ~~to various of~~ using several different physical parameterizations (Miglietta et al., 2015; Pytharoulis et al., 2018; Ragone et al., 2018; Mylonas et al., 2019). However, less attention has been ~~focused paid~~ on the microphysical processes and ~~aerosols-cloud~~ aerosol-cloud interactions, still a great source of uncertainty ~~in the understanding of~~ for understanding convective systems (Fan et al., 2016). In fact, according to the Fifth Report of the Intergovernmental Panel on Climate Change (Boucher et al., 2013), the quantification
40 of cloud and convective effects in models, and of aerosol–cloud interactions, is still a major challenge. In this type of storms, the microphysics of both cold and warm clouds ~~are essential~~ play a crucial role. The activation of aerosols as cloud condensation nuclei (CCN), the water absorption during the droplet growth, and the auto-conversion processes are main drivers in the core heating and dynamic evolution. Hence, the microphysics parameterization, along with the explicit solving of aerosols, ~~seem~~ seems to be fundamental for the development of the medicane. Gaining insight on these cloud microphysics processes is a key
45 step for reaching a complete process-understanding. In this sense, the working hypothesis in this contribution is that aerosols play a ~~part role~~ in a positive feedback with the surface winds ~~, and that fuels the storm, maintaining its structure and intensity,~~ and thus that an interactive calculation of SSA-sea salt aerosols (SSA) emissions and concentrations is fundamental for ~~the~~ development an accurate simulation of the medicane ~~, as happens~~ genesis and maintenance (Rizza et al., 2021), as in tropical cyclones (Rosenfeld et al., 2012; Jiang et al., 2019a, b, c; Luo et al., 2019).

50 In addition, ~~provided given~~ the special cyclogenetic mechanisms of medicanes, which usually need a specific atmospheric configuration to start ~~deepening~~ the convective activity, the initial conditions feeding the simulations highly affect the ~~medicane~~ development development of the medicane. In consequence, initialization time is an important source of variability (Cioni et al., 2016). In this respect, constraining the synoptic scales to follow reanalysis while allowing the model to develop the small-scale dynamics, which is exactly the function of spectral nudging, stands as a good method to reduce ~~that variability~~ this variability
55 and effectively constrain the uncertainty of the simulations.

Within this framework, ~~this article aims at analysing the present contribution aims at analyzing~~ the role played by aerosols in the development of medicanes, together with the influence of the initialisation time and the potential benefits or caveats of using nudging techniques ~~for the simulations of these storms.~~

2 Methods

60 ~~The results presented below are based on the analysis of an ensemble of 72 simulations, which consists of all the possibilities resulting from the combination of: three medicanes (Rolf, Cornelia and Celeno), two nudging configurations (no nudging -NN- and spectral nudging -SN-), two configurations for the aerosols concentration calculation (prescribed aerosols and interactively-calculated aerosols, hereinafter referred to as PA and IA, respectively) and six run-up times (12, 36, 60, 84, 108 and 132 hours).~~

65 In this section, the main techniques applied to ~~perform-conduct~~ and analyze the simulations are outlined, along with some details about the model parameterizations and the synoptic conditions associated to the studied events. It also contains a brief explanation on the interactive calculation of the SSA concentration, as included in the meteorology-chemistry coupled mesoscale model ~~WRF-chem~~ WRF-Chem.

2.1 ~~Simulations~~ Model setup

70 2.1.1 ~~Interactive vs non-interactive calculation of SSA~~

The Weather Research and Forecasting (WRF) Model (V3.9.1.1) is used to conduct the simulations object of this study (Skamarock et al., 2008). The same model configuration is employed for all simulations contained herein unless otherwise indicated. ~~No physics suite (WRF preconfiguration of a set of well-tested physics parameterizations as a suite) is used for the model run, provided the importance of choosing specific parameterizations, such as a double-moments microphysics scheme.~~ The Morrison et al. (2009) bulk ~~microphysical~~ microphysics scheme is used ($mp_physics=10$). This scheme allows for a double-moment approach, in which ~~droplet~~ number concentration, along with the mixing ratio, is considered for each hydrometeor species included in the subroutine. In its ~~single-moments version (progn=0)~~ single-moment version (progn=0), only ~~aerosols~~ the mass (i.e. mixing ratio) is taken into account. ~~Thus~~ Hence, while for the single-moment approach a constant concentration of an aerosol with a prescribed size is used, second-moment introduces the complexity degree of considering a size distribution for the aerosol population, thus being a more realistic approach.

80 ~~In the WRF-chem model, the dynamic core of WRF is coupled to a chemistry module (Grell et al., 2005). The model simulates the emission, transport, mixing, and chemical transformation of trace gases and aerosols simultaneously with the meteorology. Its main advantage with respect to using WRF alone is the possibility to perform an online calculation of the chemistry processes, which allows for chemistry-meteorology feedbacks. In the particular case under study in this paper, when using WRF alone, a fixed concentration of a given type of aerosols is prescribed everywhere. By contrast, WRF-chem calculates the aerosols distribution interactively. Specifically, the Goddard Chemistry Aerosol Radiation and Transport (GOCART) model~~

simulates major tropospheric aerosol components, including sulfate, dust, black carbon (BC), organic carbon (OC), and sea-salt aerosols, the latter being dominant in marine environments (Hoarau et al., 2018), which is our case study. GOCART includes Sea Salt Aerosol (SSA) emission as a function of the surface wind speed, initially introduced by Gong (2003) and after modified to account for SST dependence and some other corrections (Bian et al., 2019). For the emission, particles dry size is considered but the scheme also considers particles hygroscopic growth, dependent on relative humidity, according to the equilibrium parameterization by (Gerber, 1985). This dependence of SSA emission on surface wind intensity allows for the positive feedback between SSA concentration and surface wind speed that will take a main part in the medicane deepening process. When GOCART is used along with a double-moments microphysics scheme, the emission for five bulk sea salt size bins in the range of 0.06 to 20 μm in dry diameter is interactively calculated. Single-moments microphysics makes useless considering interactive aerosols.

With regard to the ~~With respect to the~~ physical configuration, radiation is parameterized with the Rapid Radiative Transfer Model for GCMs (RRTMG) by Mlawer et al. (1997), both for short and long wave radiation, solved ~~each in~~ 30 minutes ~~intervals~~. Additionally, the selected option for the surface layer ~~parameterization solves with~~ ~~parameterization uses~~ the MM5 scheme based on the similarity theory by Monin and Obukhov (1954), while the *Unified NOAA LSM* option is used ~~for to~~ ~~simulate~~ the land-surface ~~calculation interactions~~ (Mitchell, 2005). The number of soil layers in land surface model is ~~thus~~ 4. Yonsei University scheme is employed for the boundary layer (Hong et al., 2006), solved every time step (*bltdt=0*). For the cumulus physics, Grell 3D ensemble (*cu_physics=5; cudt=0*) is chosen to parameterize convection (Grell and Dévényi, 2002). Heat and moisture fluxes from the surface are activated (*isfflx=1*), as well as the cloud effect to the optical depth in radiation (*icloud=1*). Conversely, snow-cover effects are deactivated (*ifsnow=0*). ~~Landuse~~ ~~Land-use, land-sea mask topography~~ and soil category data ~~come from WPS geogrid but with dominant categories recomputed (surface input source were obtained from the WRF Users Page (WPS). SST is assimilated from ERA-Interim (sst_update=1) every 6 hours (auxinput4_interval_s = 21600). The model top is fixed at 1000 Pa and 40 vertical levels are used for the model runs.~~ Urban canopy model is not considered (*sf_urban_physics=0*), and the topographic surface wind correction from Jiménez and Dudhia (2012), ~~later modified by Lorente-Plazas et al. (2015)~~, is turned on. Both feedback from the parameterized convection to the radiation schemes and SST update (every 6 hours, coinciding with ~~boundary conditions update~~ ~~the update of boundary conditions~~) are also turned on.

2.1.1 Initial and boundary conditions

~~ERA-interim~~ ~~ERA-Interim~~ global atmospheric reanalysis is used to provide the ~~simulations with the~~ required initial and boundary conditions (every six hours). This dataset comes from ~~the~~ ECMWF's Integrated Forecast System (IFS), configured for a reduced Gaussian grid with approximately uniform 79 km spacing for surface and other grid-point fields (Berrisford et al., 2011).

All simulations are ~~performed run~~ at a 9 km of horizontal grid ~~spacing resolution~~. A different ~~non-nested~~ domain is utilized for each medicane. Domains cover the regions [16°W, 25°E, 30°N, 49°N], [4°W, 35°E, 29°N, 48°N] and [3°W, 41°E, 26°N, 45°N] in Lambert ~~conformal conic~~ projection for Rolf, Cornelia and Celeno medicanes, respectively. All the simulations ~~inside~~ ~~within~~ the ensemble of a medicane (24 ~~runs~~ per medicane) are ~~run with~~ ~~conducted in~~ the same domain.

ERA5 reanalysis dataset has been used for tracking the three medicanes with the aim of comparing the tracks obtained from the simulations with an independent source.

The specific dimensions that are changed to build up the ensemble of simulations, namely the aerosols scheme, the nudging technique and the run-up time, are described below.

125 2.1.1 Interactive versus non-interactive calculation of SSA

In the WRF-Chem model, the dynamics core of WRF is coupled to a chemistry module (Grell et al., 2005). The model simulates the emission, transport, mixing, and chemical transformation of trace gases and aerosols simultaneously with the meteorology. Its main advantage with respect to WRF alone is the possibility to perform an online calculation of the chemistry processes, allowing for chemistry-meteorology feedbacks. In the particular case under study in this contribution, when using WRF alone, a fixed concentration of a given type of aerosols is prescribed in all cells of the modeling domain. Conversely, WRF-Chem calculates the aerosol distribution interactively. Specifically, the Goddard Chemistry Aerosol Radiation and Transport (GOCART) model simulates major tropospheric aerosol components, including sulfate, dust, black carbon (BC), organic carbon (OC), and sea-salt aerosols, the latter being dominant in marine environments (Hoarau et al., 2018), as is our case study. GOCART includes SSA emission as a function of the surface wind speed, initially introduced by Gong (2003) and after modified to account for SST dependence (Bian et al., 2019). For the emission, the dry size of particles is considered but the scheme also considers the hygroscopic growth of aerosols, dependent on relative humidity, according to the equilibrium parameterization by Gerber (1985). This dependence of SSA emission on surface wind intensity allows for the positive feedback between SSA concentration and surface wind speed that plays a major role in the medicane deepening process. When GOCART is used along with a double-moment microphysics scheme, the emission for five bulk sea salt size bins in the range of 0.06 to 20 μm in dry diameter is interactively calculated. The double-moment approach (*progn=1*) has been employed to conduct all the simulations of this work. From here on, the simulations with double-moment microphysics and interactive calculation of aerosols by means of the GOCART scheme (*chem_opt=300*) will be referred to as "IA" simulations, while for those with double-moment microphysics but prescribed aerosol concentration (*chem_opt=0*) the term "PA" will be used.

145 2.1.2 Spectral nudging

Spectral nudging is a technique for constraining the synoptic scales to follow reanalysis while allowing the regional model to develop the small-scale dynamics (Miguez-Macho et al., 2004). Initially conceived for reducing the sensitivity of regional climate simulations to the size and position of the domain chosen for calculations, it has been suggested that this technique is necessary for all downscaling studies with regional models with domain sizes of a few thousand kilometers and larger embedded in global models, in order to avoid the distortion of the large-scale circulation. With this premise, we analyze the effects of considering ~~the introduction of~~ spectral nudging for the simulation of medicanes. Particularly, a wavelength of 999 km in both horizontal directions has been used to ensure that only synoptic scale dynamics are constrained; wind, temperature and water ~~vapour~~-vapor mixing ratio fields are nudged above the PBLplanetary boundary layer.

2.1.3 Run-up time

Run-up time makes reference to the time period (in hours) since the start of a simulation (reference time) to the time in which the medicane appears. To follow a consistent criterion, this reference time is extracted from the complete ensemble of each medicane. For example, for Celeno medicane, the start reference time is considered to be 1995 January 14 12:00 GMT, as can be extracted from Figure 1 on the main text UTC (Figure 1). Hence, six different initialization times are considered for the ensemble of Celeno medicane simulations ensemble: from 1995 January 09 00:00 GMT-UTC to 1995 January 14 00:00 GMT-UTC with 1 day intervals, corresponding to 132, 108, 84, 60, 36 and 12 hours of run-up time, respectively. The same six run-up times are considered for Rolf (reference time 2011 November 06 12:00 GMT-UTC) and Cornelia (reference time 1996 October 06 12:00 GMT-UTC) medicanes. By considering an ensemble of initialization times, in practice we are perturbing we are in practice changing the initial conditions to reduce-constrain the possible uncertainty associated to this factor, thus producing more consistent results for the sensitivity of using or not when addressing the sensitivity to using (or not) an interactive aerosols calculation.

2.2 Synoptic environments of the events

The Rolf medicane, also known as Tropical Storm Rolf, Tropical Storm 01M (NOAA National Oceanic and Atmospheric Administration -NOAA-) or Invest 99L (NRL U.S. Naval Research Laboratory) was a Mediterranean tropical-tropical-like storm occurred on 2011 November 06-09. It started from the alignment of a surface low-pressure system which evolved into a baroclinic environment near the Balearic Islands, and an extensive upper-level trough early on November 06. On November 07, the system revealed tropical characteristics such as a warm core and convective bands organized around a quasi-symmetric structure. Early on November 08, Rolf reached its maximum intensity, and NOAA officially declared the system a tropical storm (Win). Reaching its peak intensity on that same day (991-990 hPa of central pressure and maximum 1-minute sustained winds of 82 km/h), Rolf started to weaken, transitioned to a tropical depression and finally lost its structure late on November 09 when it made landfall in Southeast France (Dafis et al., 2018; Miglietta et al., 2013) (Ricchi et al., 2017; Dafis et al., 2018; Miglietta et al., 2013). Rolf was the first tropical-tropical-like cyclone ever to be officially monitored by the NOAA in the Mediterranean Sea.

Known to have formed from the interaction of a large compact low-pressure area that approached to Greece from the Ionian Sea, and a middle tropospheric trough that extended from Russia to the Mediterranean, Celeno started its convective activity early 1995 January 14. Initially remaining stationary between Greece and Sicily with a minimum atmospheric pressure of 1002 mb at hPa, the newly formed system began to drift southwest-to-south in the following days influenced by northeasterly flow incited by the initial low. It acquired a cloud-free distinct eye and a spiralling rainband, rapidly deepening and started a rapid deepening phase. Its track is generally depicted crossing the Ionian Sea southwards, from Southern Greece to the coast of Lybia (Pytharoulis et al., 2000). A strong disagreement seems to exist between the different sources of information about the maximum wind speed and minimum central SLP reached by Celeno tropical-like cyclone. The German research ship Meteor noted winds of 135 km/h, and SLP of 980 hPa was registered in the North Coast of Malta. ERA5 reanalysis provides an SLP

~~under of~~ 990 hPa on 1995 January 14 ~~1012:00 GMT-UTC~~ for this cyclonic structure. According to Lagouvardos et al. (1999), a ship near to the vortex centre (35.7°N - 18.2°E) reported a 83 km/h surface wind and a pressure of 1009 hPa on 1995 January 16 06:00.

190 ~~Lastly,~~ Last, the first phase of Cornelia medicane took place between the Balearic Islands and Sardinia, with an eye-like feature clearly developed. It appeared on 1996 October 06 to the north of Algeria, and strengthened before temporarily losing its ~~structure~~ eye-like structure when making landfall in Sardinia. On October 09, the system strengthened again over the Tyrrhenian Sea and passed north of Sicily, reporting winds of ~~90-81~~ km/h at 100 km from the storm ~~'s center-center~~ (a ship in the position 40°N - 13°E). The lowest ~~model~~ estimated atmospheric pressure reached in the storm center was ~~998 hPa~~ (Pytharoulis et al., 2000; Cioni, 2014; Caviechia and Von Storch, 2011)996 hPa (Reale and Atlas, 01 Feb. 2001; Pytharoulis et al., 2000; C

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2.3 Methods for the analysis of the simulations output

2.3.1 Tracking algorithm

TITAM (Pravia-Sarabia et al., 2020) is an algorithm specifically suited to allow for the detection and tracking of medicanes even in adverse conditions, such as the existence of a large low in the domain, or the coexistence of multiple medicane 200 structures. This algorithm, based on a time-independent approach, has been used to study the intensity and duration of the medicanes reproduced in the different simulations presented in this ~~paper. Given the different nature of each medicane in duration, intensity and spatial extent, the namelists parameters for TITAM must be consistently defined prior to its execution on each medicane .~~ Specifically, the 4th Hart condition (thermal wind magnitude greater for the lower than for the upper tropospheric layer) ~~is not checked for medicane tracking in Celeno simulations~~ contribution. For the tracking of medicanes in

205 both the WRF simulations and the ERA5 reanalysis dataset, the following parameters have been chosen for running TITAM algorithm: 5 smoothing passes for the cyclonic potential field; 1020 hPa of SLP threshold (no structure is discarded by the SLP filter); 0.5 rad·h⁻¹ as vorticity lower threshold; a zero vorticity radius required to be symmetric in 4 directions with lower and upper thresholds of 50 and 500 km, respectively, with a maximum allowed asymmetry of 300 km; 5 minimum points in a cluster to be considered a medicane structure; a symmetry Hart parameter (B) calculated in four directions in the 900-600

210 hPa layer, having the maximum B an upper threshold of 20 m; and the lower and upper tropospheric thermal wind parameters calculated in the 900-600 and 600-300 hPa layers, respectively. The zero vorticity radius, which is time- and point-dependent, is the radius used to calculate the Hart parameters for each medicane center candidate on each time step.

2.3.2 ~~Medicanes~~ Medicane duration ~~and intensity~~

~~When talking about duration, two different measures will be distinguished. Duration will be~~ Duration is associated to the 215 number of ~~points~~ model time steps in which the algorithm detects a medicane. ~~By contrast, support will be~~ However, instead of using the total length, the duration of the medicane is calculated as the ~~more compact points set~~ most compact set of points. This compact set, as shown in Figure ~~1 on the main text 1~~ (grey boxes), ~~serve~~ serves as an objective measure of the real duration

of the medicane, removing early starts and late endings in which ~~medicane-structure~~ the structure of the medicane is not well defined, ~~repeatedly-noisily~~ gaining and losing medicane condition. To calculate ~~this-the~~ compact set for a given simulation, after having run the tracking algorithm ~~on-it,~~ take a vector X , it takes a vector X in which elements x_i , each one corresponding to ~~a-time-step,~~ an output step (e.g., hourly), are 1 if a medicane is found for the ~~time-output~~ step, and 0 if not. For each i in ~~$1:N_t-1$~~ $1:(N_t-1)$, and each j in ~~$i:(N_t-1)$~~ $(i+1):N_t$, find pair ~~$[i,j]$~~ $[i,j]$ such that:

$$Q_{i,j} = \sum_{m=i+1}^{m=j} x_m + \sum_{m=i+1}^{m=j} (x_m - 1) \quad (1)$$

is maximum, being ~~N_t~~ N_t the number of ~~time-output~~ steps in the simulation. Once ~~found~~ the pair ~~$[i_m, j_m]$~~ that makes this ~~quantity-making~~ $Q_{i,j}$ maximum is found, ~~i_m and j_m~~ i_m and j_m are respectively the initial and final ~~positions-time steps~~ positions-time steps of the medicane ~~more compact set,~~ and thus most compact set of points, and its difference is ~~an-objective-measure-of-the-medicane-duration~~ used hereinafter as the measure for the duration of the medicanes.

2.3.3 Kernel density estimation

In order to represent the density of medicane positions over space, Kernel Density Estimation (KDE) is employed. Let ~~p_i~~ p_i be ~~the position of a medicane at a given time,~~ defined by a pair (x_i, y_i) of longitudinal and latitudinal positions in the matrix. Then, the density estimate for a set of medicane positions is calculated as:

$$f_H(\mathbf{x}) = \frac{1}{n} \sum_{i=1}^n \mathcal{K}_H(\mathbf{x} - p_i) \quad (2)$$

~~where H is the bandwidth and n the number of points in the sample from which the density estimate is drawn.~~ For the cases contemplated throughout this contribution, a Gaussian Kernel with a bandwidth of 2 grid points is used.

235 3 Results

The results presented below are based on the analysis of an ensemble of 72 simulations, which consists of all the possibilities resulting from the combination of: ~~three medicanes (Rolf, Cornelia and Celeno),~~ two nudging configurations (no nudging and spectral nudging), ~~two configurations for the aerosols concentration calculation (interactively-calculated aerosols and prescribed aerosols, hereinafter referred to as IA and PA, respectively)~~ and six run-up times (12, 36, 60, 84, 108 and 132 hours).

3.1 Analysis of the ensemble of simulations

As an initial ~~exampleapproach,~~ a detailed view of each ensemble member for the Celeno medicane (the most intense out of the three simulated) during its lifetime when no nudging is applied is ~~three medicanes during their lifetime is~~ shown in Figure 1.

Note that although the discussion below is mostly based on this case, analogous figures for the rest of cases are presented in Section S1.1 of Supplement, and the coherence across cases, as well as relevant differences, are commented here. For each run-up time, the pair PA/IA is depicted; a dot-circle represents a time step where a medicane is found, being its colour the sea-level pressure (SLP) value for the medicane center. If a medicane is not found for a given time step, a grey cross is placed. Figure 1 shows how, when no nudging is used (left column plots), longer and deeper storms are systematically reproduced for IA simulations with respect to the PA simulations. The application of spectral nudging (see Figure S5 in Supplement right column plots) makes IA produce weaker medicanes than for the no nudging case, but the medicanes reproduced by IA are still longer and deeper in SLP than for PA. The role of initialization time is also clearly depicted in this figure Figure 1: it induces a noticeable but nearly random behavior on the medicane response, with differences up to 5 hPa on the center-central SLP of the medicanes for two consecutive run-up times (i.e., initialization times separated 1 day), but in opposite directions both directions and without a discernible pattern. Similar results of longer and deeper tracks for IA, the However, spectral nudging reduces this variability introduced by initialization time and the spectral nudging reduction of the IA influence are found for the other two storms (associated plots are included in Section S1.1 of Supplement). the initial conditions, evening out the runs output, and sometimes even producing longer-lasting medicanes (e.g., for the case of Rolf). These results are consistently reproduced for the three medicane cases considered.

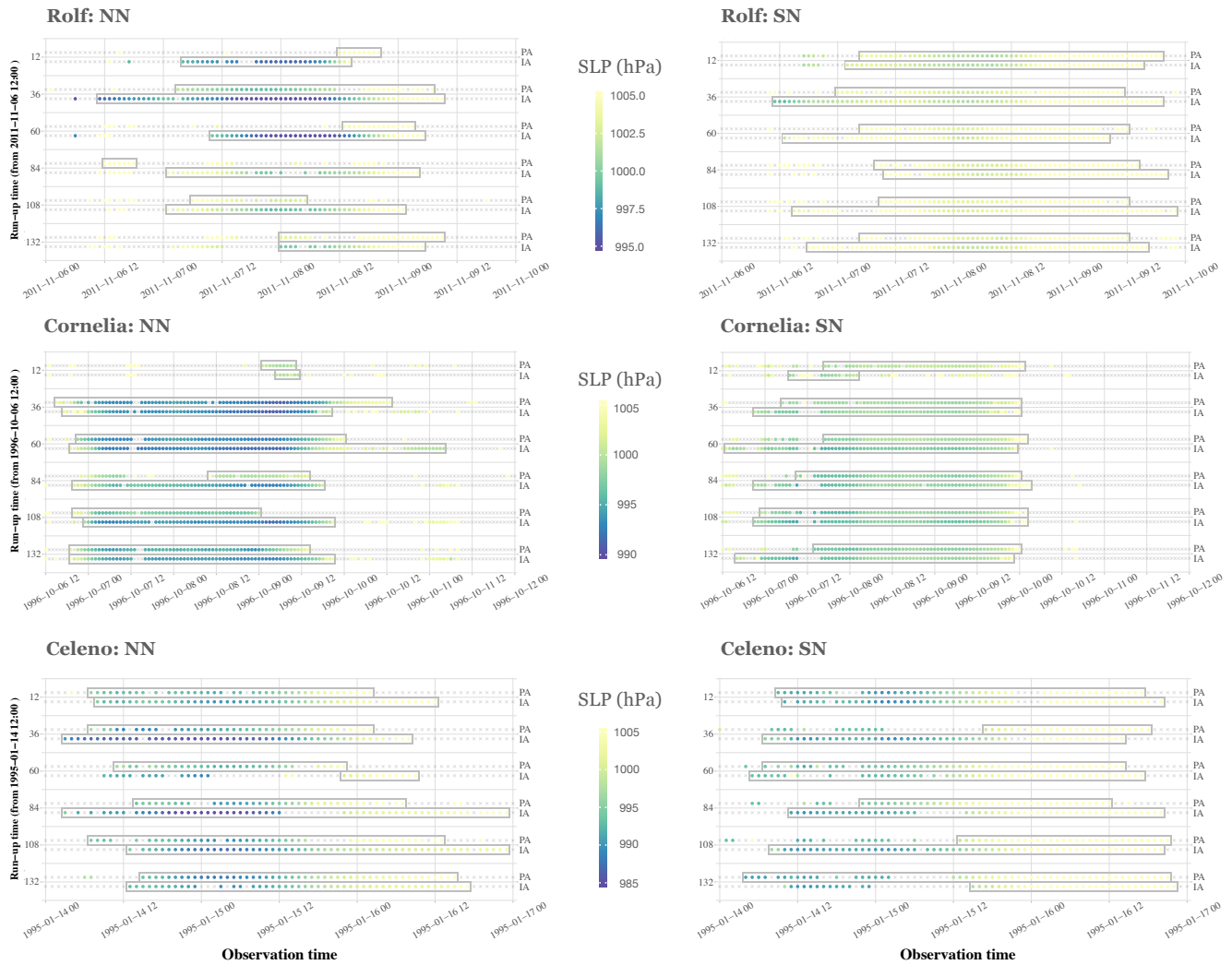


Figure 1. PA/IA pair-pairs of simulations without (left) and with spectral nudging (right) represented for each run-up time of initialization for medicane-medicane Rolf (top), Cornelia (middle) and Celeno (bottom). A dot-circle represents a time step where a medicane is found, being its colour the SLP value for the medicane center. A grey cross is placed for the time steps in which no medicane is found. The grey frames include, for each simulation, the time steps inside the medicane more compact points set of points, as described in Subsection 2.3.2.

In order to study not only the track duration, but also their spatial differences, Figure 2 illustrates the trajectories of the different simulations without nudging for Cornelia medicane (the longest-lasting event of the three considered). The sensitivity of the medicane tracks for the different simulations is illustrated in Figure 2. The data is aggregated across different the run-up times through the calculation of a Kernel Density Estimation (KDE) from most likely KDE with Gaussian kernel from most-likely cyclone locations (Rosenblatt, 1956) built on top of the medicane-center-center of the medicane positions along the tracks belonging to the PA (green) and IA (red) ensembles of simulations without nudging. KDEs for all cases can be found in

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265 ~~Section S1.2 of Supplement. A threshold of 0.2 is used to convert KDEs, different ensembles of simulations. For each medicane (Rolf, left; Cornelia, center; Celeno, right), the four nudging-aerosols (NN vs SN; PA vs IA) simulations ensembles, each one with six run-up times, are converted into a KDE, normalized to the [0,1] range, to binary density clouds as those included in Figure 2. In addition to how the storm severity is increased for the IA ensemble (as previously shown), the figure showeases that their-. The three medicanes tracks obtained as a result of running the TITAM tracking algorithm (Pravia-Sarabia et al., 2020)~~
270 ~~on ERA5 reanalysis data is superimposed for the sake of comparison. Figure 2 depicts that the tracks are more spatially constrained and longer when IA is considered for the IA ensemble. With the introduction of spectral nudging(see Figure S9 of Supplement), tracks are spatially constrained, thus limiting the intensification potential of the medicane. Similar results can be drawn from the analysis of the Rolf and Celeno medicanes , as extracted from Figures S7, S8, S11 and S12 of Supplement~~
275 ~~Rolf tracks turn stationary and do not reproduce the observed movement of the storm, while a better agreement with the ERA5 track is achieved for Cornelia, specially in the latest phase, when the storm was moving towards Sicily. The case of Celeno is an extreme case in which the spectral nudging is needed for the simulations to replicate the real track (although a medicane does develop when no nudging is considered). Hence, IA produces deeper and longer medicanes with a more spatially constrained trajectory; and spectral nudging seems to be beneficial or detrimental depending on the case without a clear pattern.~~

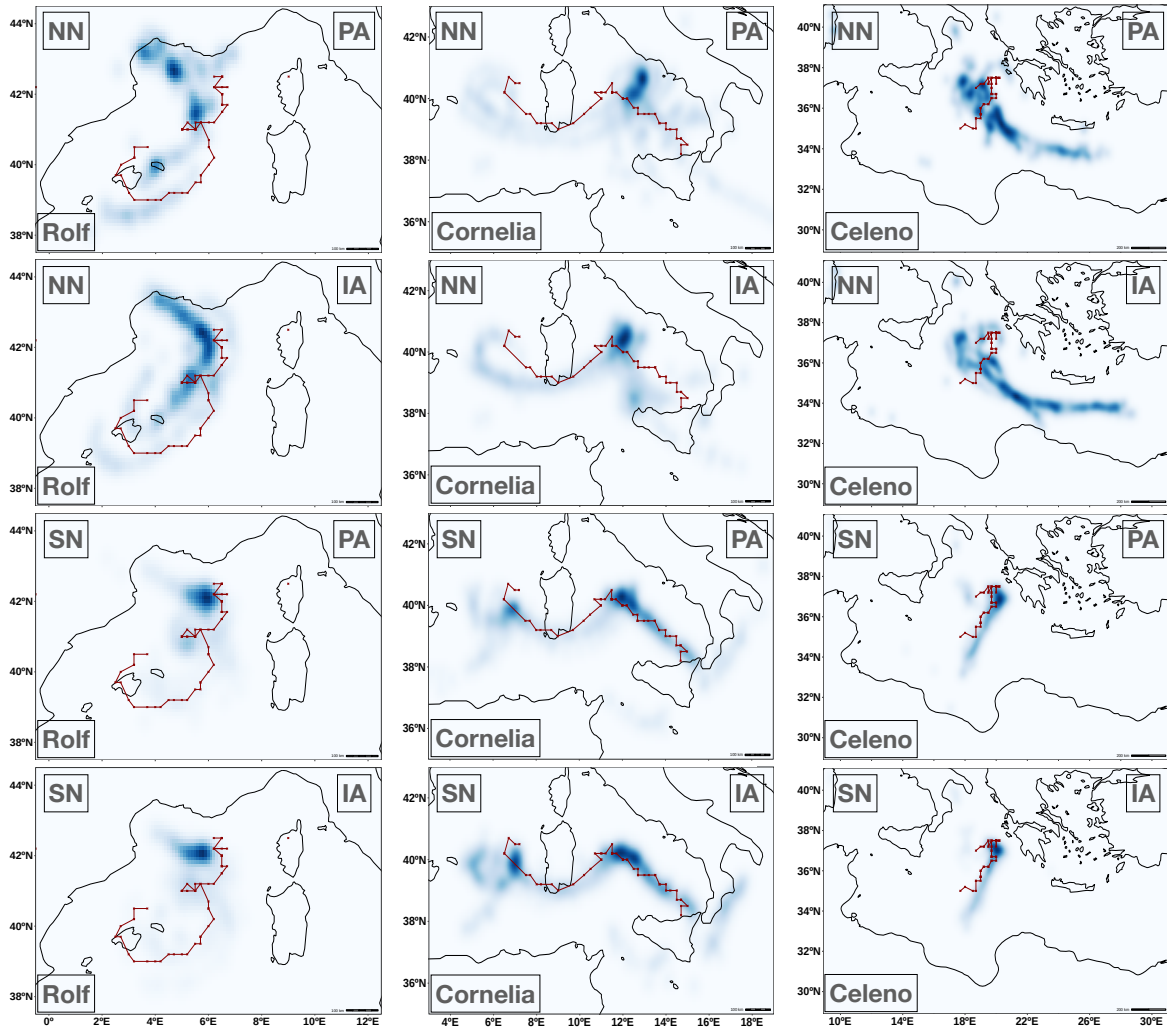


Figure 2. Figure shows the binary-Normalized KDEs built on top of the medicane positions along the tracks belonging to the PA nudging-aerosols ensembles for medicanes Rolf (greenleft), Cornelia (center) and IA-Celeno (redright)simulations without nudging ensembles for. In red, the track of each medicane Corneliaas a result of running TITAM (Pravia-Sarabia et al., 2020) tracking algorithm on ERA5 reanalysis data is superimposed.

To offer a complete-more comprehensive view of the results, Figure 3 summarizes the main outcome of the simulations ensemble, focusing both on the differences in for each member within the ensemble of simulations. Medicanes are separated in rings; and colours indicate, for the first two rows, the minimum SLP reached at the center of the medicane during its lifetime and its total duration. Medicanes are separated in rings and colours indicate in each simulation for the four nudging-aerosols ensembles. The widths of the ring sector are proportional to the relative duration of the events (with respect to the maximum

duration of a simulation inside the event simulations ensemble). The two bottom panels focus on the difference in depth of the storm when ~~SSA-interactive calculation~~ the interactive calculation of aerosols is considered or not (reddish colours indicate deeper storms when IA is considered), ~~while ring-sector widths~~ thus being the third row the result of subtracting the first from the second row. For these two panels, widths of the outer rings are proportional to the ~~relative duration of the events~~ (with respect to the maximum duration of a simulation inside the event simulations ensemble). Both panels are based on identical ~~calculations~~, but length of the most compact set of points for the IA simulations, and widths of the inner rings to that length for the PA ones. Differences between left and right panels illustrate the impact of using spectral nudging (right) versus leaving simulation free (left). ~~Differences are clearly positive for most cases, which involves that, in general, IA leads to more intense medicanes. Moreover, medicane tracks are also longer for IA simulations. Spectral nudging does not change this general conclusion, but reduces the effect of introducing IA, and generally leads to shorter medicane tracks~~In line with what was concluded from previous figures, this Figure 3 summarizes and highlights the fact that in NN simulations, IA produces deeper and longer-lasting medicanes as compared to those reproduced with PA. With respect to the initialization time, there seems to be a ~~nearly~~ random response to the initial conditions ~~-(variability in the azimuthal direction)~~. The use of SN drastically ~~-yet not completely-~~ reduces these differences but leads to even longer-lived medicane structures (which, for the cases of Cornelia and Celeno, approaches the simulations to the observed tracks), and reduces the variability introduced by the run-up time.

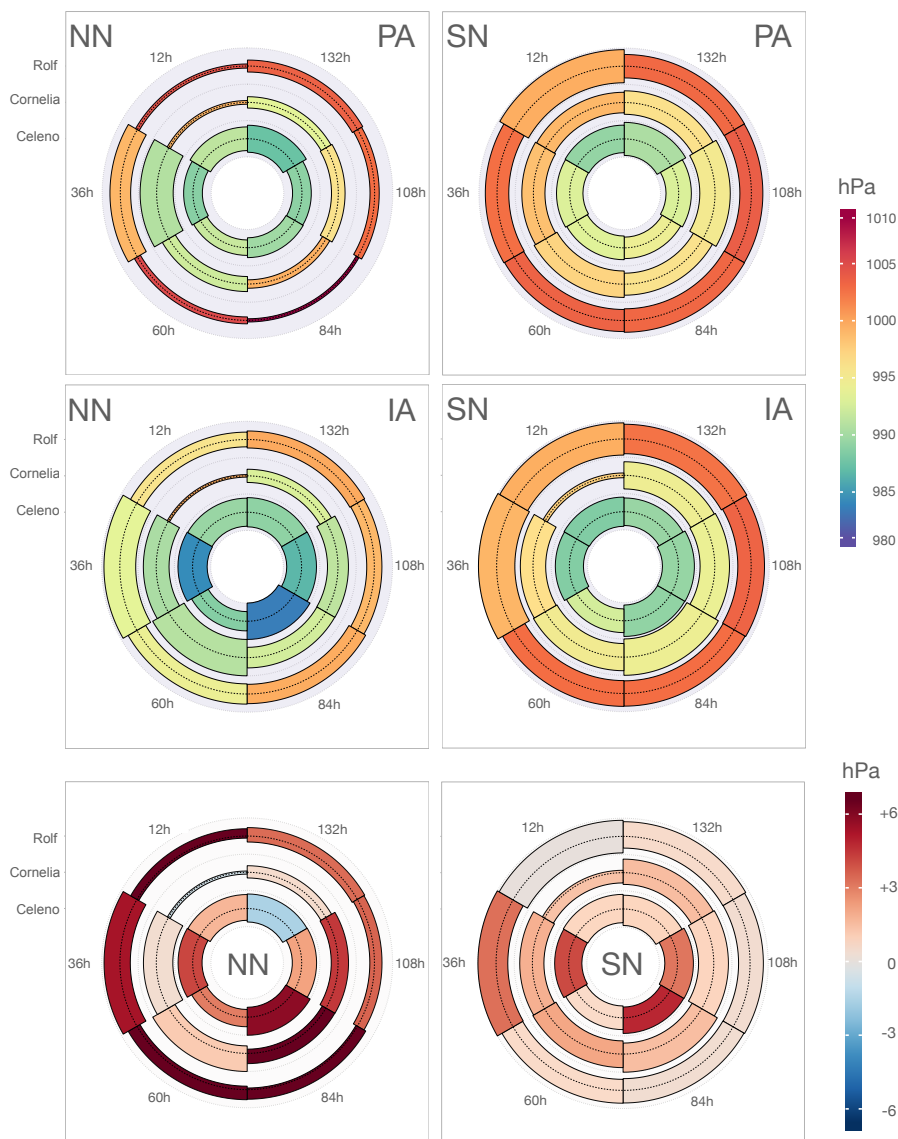


Figure 3. Absolute difference (colours) in the minimum SLP among the medicane centers reached during the medicane lifetime between the IA and PA simulations, for each medicane and run-up time. Medicanes are separated into rings across radial direction (inner Celeno, middle Cornelia, outer Rolf). Different The different run-up times are shown in the six ring portions. For In the first two rows, colours indicate the minimum SLP reached in each ring portion simulation, for the width of the upper half is four nudging-aerosols simulations ensembles; ring sector widths are proportional to the number relative duration of points in which a medicane is found in the IA simulation events. In the last row, and that of the lower part to the number of points absolute difference (colours) in the PA simulation. A dashed line separates both. Widths of each minimum SLP among the medicane are durations relative to centers reached during the maximum duration among all medicane lifetime between the IA and PA simulations is shown, for the each medicane (24 simulations) and run-up time. Thus, a red colour represents a deeper SLP minimum of the medicane centers SLP values for the IA simulations, and the more larger the asymmetry between the widths of the upper and lower halves of each portion of the rings, the higher the influence of using IA in the track medicane duration. For each ring portion, the width of the outer half is proportional to the length of the compact set of points in the IA simulation, and that of the inner part to the one in the PA simulation. A dashed line separates both. All rings widths are normalized with respect to the maximum duration among all the simulations ensemble for the medicane (24 simulations).

4 Proposed intensification mechanism: SSA-wind feedback

300 A more quantitative analysis is shown in Tables 1 and 2, containing both the minimum SLP reached by the medicane center and the duration of the medicane (length of the compact set of points) reproduced in each simulation, respectively. The μ quantity reduces the ensemble through the sample mean for each aggregated recursive level (aerosols -ensemble with the different run-up times-; nudging -ensemble with the different run-up times and PA/IA configurations-; medicane -ensemble with the different run-up times, PA/IA configurations and NN/SN configurations-).

Table 1. Summary of the minimum SLP reached by the medicanes in the 72 simulations ensemble. The μ quantity reduces the ensemble through the sample mean for each aggregated recursive level.

Medicane →	Rolf				Cornelia				Celeno			
Nudging →	NN		SN		NN		SN		NN		SN	
Aerosols →	PA	IA	PA	IA	PA	IA	PA	IA	PA	IA	PA	IA
Run-up time ↓												
12h	1003.8	995.8	999.3	999.3	998.9	999.5	998.7	997.3	991.6	989.9	989.7	988.9
36h	998.8	993.5	1002.1	998.9	991.1	990.7	998.3	996.4	989.2	985.1	993.1	989.0
60h	1004.3	994.4	1002.8	1002.2	992.4	991.3	997.0	995.0	992.1	989.1	993.3	992.6
84h	1012.6	999.4	1002.6	1002.2	999.2	992.6	996.0	994.5	990.2	984.4	994.4	989.6
108h	1002.3	998.7	1003.1	1002.7	996.0	991.6	995.1	994.2	989.5	987.2	993.0	989.8
132h	1002.9	999.5	1002.5	1001.9	993.5	993.0	996.1	994.6	988.0	989.6	990.8	989.9
μ (aerosols)	1004.1	996.9	1002.1	1001.2	995.2	993.1	996.9	995.3	990.1	987.6	992.4	990.0
μ (nudging)	1000.5		1001.6		994.1		996.1		988.8		991.2	
μ (medicane)	1001.1				995.1				990.0			

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Table 2. Summary of the medicane duration -number of points in the compact set- of the 72 simulations ensemble. The μ quantity reduces the ensemble through the sample mean for each aggregated recursive level.

Medicane →	Rolf				Cornelia				Celeno			
Nudging →	NN		SN		NN		SN		NN		SN	
Aerosols →	PA	IA	PA	IA	PA	IA	PA	IA	PA	IA	PA	IA
Run-up time ↓												
12h	9	32	72	70	10	7	56	13	43	51	50	53
36h	45	69	51	79	90	70	63	69	34	53	48	51
60h	15	44	48	56	41	99	72	74	28	35	43	43
84h	7	42	51	55	23	56	60	99	36	66	39	57
108h	20	33	50	58	37	57	83	80	34	54	46	58
132h	26	37	51	64	31	36	64	74	49	53	60	50
μ (aerosols)	20.3	42.8	53.8	63.7	38.7	54.2	66.3	68.2	37.3	52.0	47.7	52.0
μ (nudging)	31.6		58.8		46.4		67.3		44.7		49.8	
μ (medicane)	45.2				56.8				47.3			

310 With respect to the spectral nudging effect, although long-lived medicane structures are generated for SN configuration, they do not reach the intensity of those reproduced in the NN simulations. The explanation to this effect lies in the spectral nudging mechanism: forcing the meteorological fields to resemble to large scale dynamics produces alterations in the nudged fields (temperature, humidity and wind) above the PBL. Henceforth, the temperature field does not freely evolve and deep convection may be interrupted, thus limiting the intensification potential of the medicane. Figure 4 supports this statement, showing that the warm core is broken-off in the 500-800 hPa layer when SN is introduced. In Figure 4, a set of height-radius cross sections of the equivalent potential temperature (θ_e) is produced by time-averaging the cross-sections for the time steps in which a medicane is found for the NN (top) and SN (bottom) simulations of medicanes Rolf, Cornelia and Celeno, started

315 with 36 hours of run-up time.

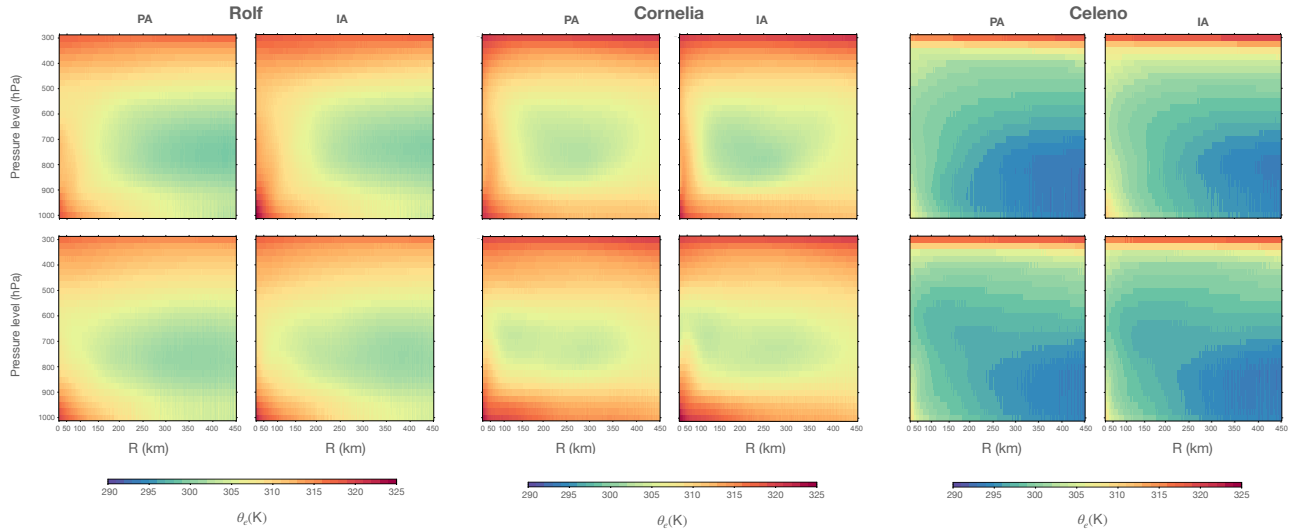


Figure 4. Time-averaged height-radius cross sections of the equivalent potential temperature (θ_e) over all time steps in which a medicane is found in the NN (top) and SN (bottom) simulations for medicanes Rolf (simulation starting 2011 Nov 05 00:00 UTC -36 hours of run-up time-), Cornelia (simulation starting on 1996 Oct 05 00:00 UTC -36 hours of run-up time-) and Celeno (simulation starting on 1995 Jan 09 00:00 UTC -36 hours of run-up time-).

3.1 Proposed intensification mechanism: SSA-wind feedback

As aforementioned, it seems clear that It has been previously discussed that IA calculation leads to deeper and longer medicane tracks. As introduced in Section 1, our main initial hypothesis, provided-given the close nature of medicanes to tropical cyclones, is that the explicit-online calculation of SSA in-the-model-aerosol-scheme-(GOCART)-allows for the existence of a positive feedback with surface wind. Although this feedback is irrelevant for an early emergence of convective activity, generally fostered by a cold cut-off low in upper levels when it comes to medicanes (Emanuel, 2005), it becomes essential once the core circulation is established. For the sake of examining the hypothesis of the existence of this feedback, we choose two simulations-of-medicane-Rolf-without-nudging-NN simulations of Rolf starting on 2011 November 05 00:00 GMT. Figure 4 UTC, because (1) their closeness to the observed medicane track; (2) the low SLP they reach; and (3) the robust and stable structure they develop (with the exception of the intensity, it is the IA/PA pair in which a most similar storm is developed in both simulations). Figure 5 depicts the temporal evolution of the differences in equivalent potential temperature (θ_e) between the IA and PA simulations along the vertical of the medicane center.

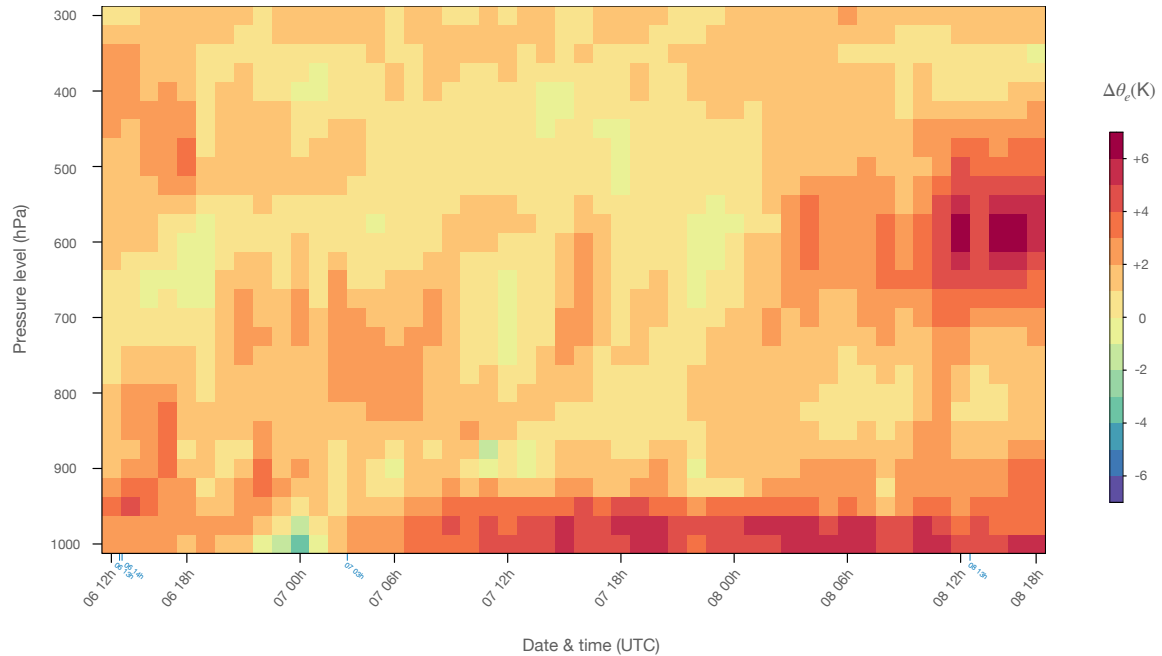


Figure 5. Difference in equivalent potential temperature between the IA and PA simulations for each output time step (horizontal axis) in which both the PA and IA simulations find a medicane for Rolf starting on 2011 Nov 05 00:00 UTC. Results are presented for each pressure level (vertical axis) in the vertical of the center of the medicane. Red colours indicate higher equivalent potential temperatures, and thus more available energy, for the IA simulation. Blue marks in the horizontal axis refer to the time steps in which the medicane structure is lost and hence the tracking algorithm does not find a medicane center. For these time steps, there is no vertical profile of temperature of the medicane core. The time axis labels (horizontal axis) are presented in day-hour format, being all dates referred to November 2011.

In general, higher (orange to red colours) θ_e values are found for the IA case, ~~which means that higher~~ meaning that more convective potential energy comes into play in the form of hot moist air ~~presence~~ in the low troposphere when aerosols are interactively solved in each ~~timestep~~ time step of the model. This is specially ~~true for timesteps later than~~ evident for time ~~steps after~~ steps after November 07, when ~~the~~ the medicane starts to gain strength, ~~releasing more latent heat of condensation due to the~~ stronger convection, and thus producing more water vapour under higher surface wind conditions. Another related aspect is ~~shown in Figure 5~~, the warm core structure, is presented in Figure 6 by means of the ~~differences between the IA and PA~~ time-averaged ~~height-longitude~~ height-radius cross sections of θ_e ~~It~~ (top plots) along the latitudes of the medicane center, for
 335 both the PA (left) and IA (right) NN simulations of Rolf started with 36 hours of run-up time. Figure 6 clearly reveals that the core ~~temperature is up to 4 degrees~~ equivalent potential temperature is higher for the IA case, ~~and even the divergence aloft can~~ be appreciated in this figure ~~this being specially true for the lower tropospheric and surface levels given the strong presence of~~ moist air in this layer; these plots even depict the divergence of warm air aloft. In the bottom plots of Figure 6, the anomaly of potential temperature (θ') in time-averaged height-radius cross sections along the latitudes of the medicane center is presented

340 for the same two simulations, showing a net heating of the storm core, maximum in the 500-800 hPa layer, which seems to be related to core dynamics.

For each time step (horizontal axis) in which both the PA and IA simulations find a medicane for Rolf medicane starting on 2011-Nov-05 00:00 GMT, the difference in equivalent potential temperature (in the medicane center point) between the IA and PA simulations is presented for each pressure level (vertical axis). Red colours indicate higher equivalent potential temperatures, and thus more available energy, for the IA simulation. Blue marks in the horizontal axis refer to the time steps in which the medicane structure is lost and, therefore, the tracking algorithm does not find a medicane center.

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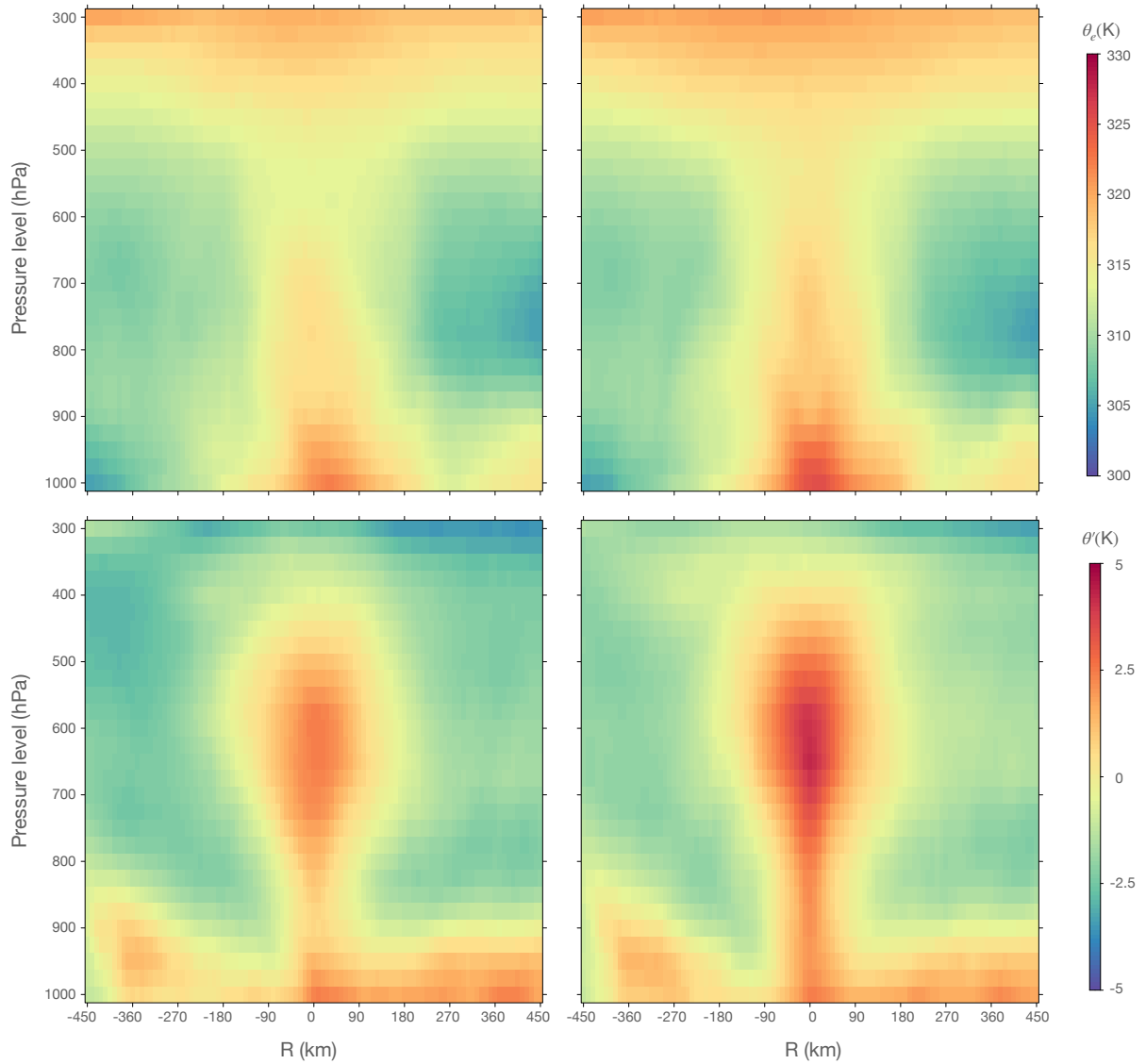


Figure 6. Figure shows differences between the IA and PA. Top row: time-averaged height-longitude cross sections of θ_e along the medicane center latitudes during, for both the medicane lifetime. The simulations are the PA (left) and IA (right) NN simulations of Rolf medicane cases without spectral nudging starting on 2011 November 05 00:00 GMT started with 36 hours of run-up time. Distance from Bottom row: for the medicane center is represented in the horizontal axis same two simulations, and atmospheric pressure levels the in-level anomaly of potential temperature (θ') in time-averaged height-radius cross sections along the vertical axis medicane center latitudes is presented.

To fully understand the processes undergone by aerosols and clouds, the ~~hydrometeors distribution~~ distribution of hydrometeors is examined, revealing the form in which the thermal energy is handled by the system. This is shown in Figure 67, in which a ~~time-averaged time- and azimuthally-averaged~~ height-longitude cross section is presented along the ~~latitudes of the~~ med-
350 icane centers found for the simulations ~~without nudging of Rolf~~ starting on 2011 November 05 00:00 GMT. ~~In this plane,~~
~~the following projections are presented~~ UTC (36 hours of run-up time) with prescribed (left) and interactive (right) aerosols
calculation, both simulations run without spectral nudging. 7 shows (from top to bottom) ~~cloud water~~ mixing ratio, ~~droplet~~
~~number~~ rain water mixing ratio and ~~rain water~~ droplet number mixing ratio, for both the PA (left column) and IA (right column)
simulations.

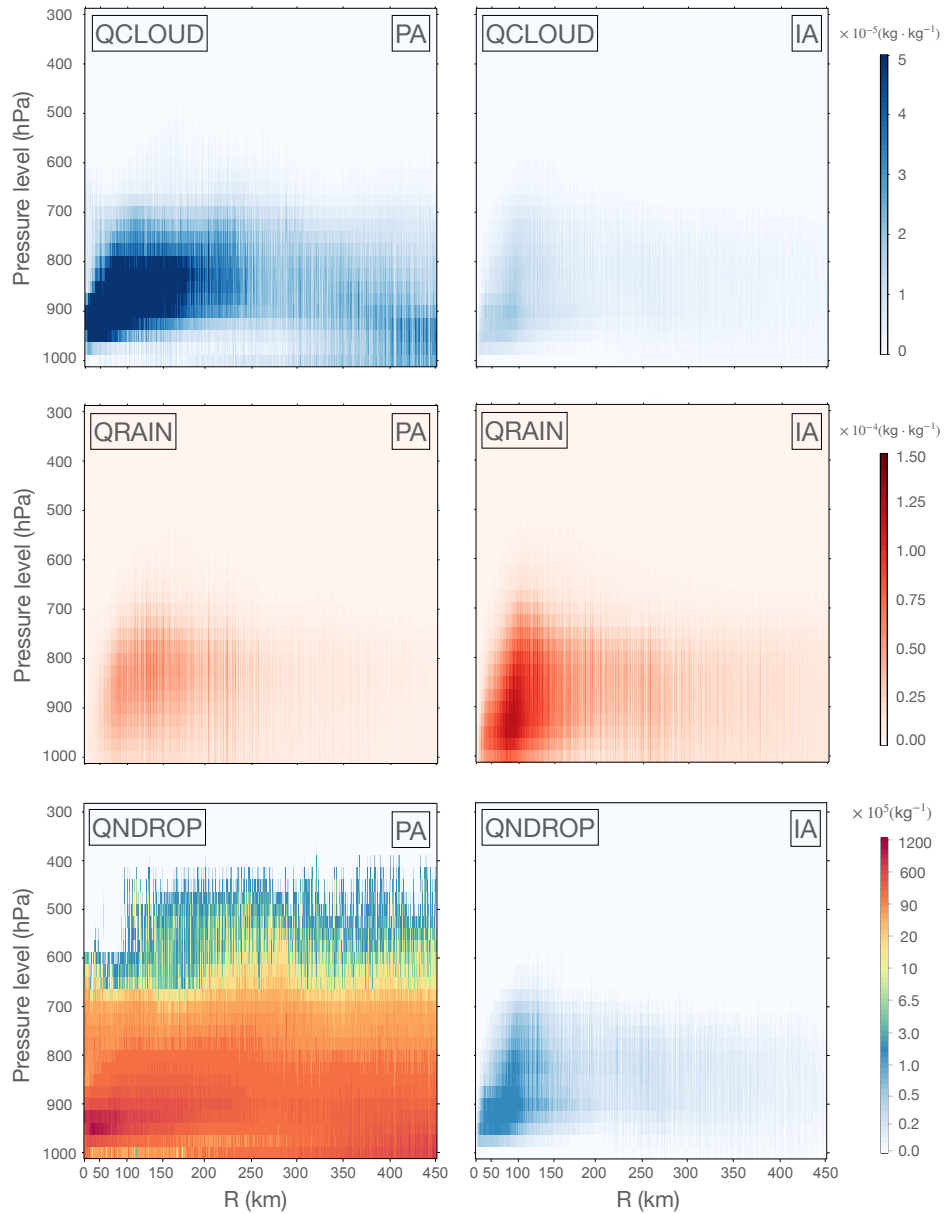


Figure 7. From top to bottom: cloud water mixing ratio (Q_CLOUD, kg kg^{-1}), droplet number rain water mixing ratio (Q_RAIN, kg kg^{-1}) and rain water droplet number mixing ratio (Q_DROP, kg kg^{-1}). For the three fields, a time-averaged height-longitude-Time- and azimuthally-averaged height-radius cross section along the medicane center latitudes during its lifetime sections are shown presented, for PA (left column) and IA (right column) simulations of Rolf starting with 36 hours of run-up time (i.e., on 2011 November 05 00:00 GMTUTC), both without spectral nudging. Distance from the medicane center is represented in the horizontal axis r , and atmospheric pressure levels in the vertical axis. Logarithmic scale is used for The three scales must be multiplied by the droplet number mixing ratio factor preceding the units in the legends titles.

355 Figure 6.7 indicates that in the case where the largest medicane deepening is found, although ~~being~~ accompanied by a higher thermal energy, lower cloud droplet numbers and less cloud water content come into play, but more rain water is produced. Diving into the microphysics, a plausible explanation resides in the Köhler curves and mechanism of CCN activation as cloud droplets. Intense wind blowing over the ocean surface creates sea spray, which contains organic matter and inorganic salts that form SSA (Gong et al., 1997). SSA are coarse particles quickly reaching the critical radius and being early activated as CCN at
360 low supersaturation rates according to Köhler theory (Köhler, 1936), thus being highly prone to condensational growth (Jensen and Nugent, 2017). The existence of SSA enables an early, rapid and strong latent heat release in the lower troposphere, which enhances deep convection, ultimately leading to an intensification of surface winds under low supersaturation conditions in the early medicane stage. Conversely, prescribed aerosols concentrations used in PA simulations lead to a high amount of fine particles with low hygroscopicity hardly activated as CCN and competing for water ~~vapour~~-~~vapor~~ uptake, thus producing a
365 higher number of small droplets which are barely converted into raindrops.

4 Summary and conclusions

In this contribution, an ensemble of ~~simulations have~~ 72 simulations has been conducted to analyze the role of SSA feedbacks in the development and intensification of three different medicanes. Results show a clear dependence ~~on interactive aerosols calculation~~ of both the track and intensity of the medicanes simulated on the calculation of interactive aerosols, as
370 their consideration leads to ~~reproducing~~ longer and deeper medicanes in the simulations. The proposed mechanism to explain this difference is that, by contrast with simulations with prescribed aerosols (as usually included in meteorological models), when interactive aerosols are introduced the presence of coarse particles is ~~accounted~~ taken into account, and the hygroscopic characteristics of ~~sea salt aerosols~~ SSA are considered, thus allowing an early reaching of critical radius and an enhancement of a strong latent heat release in the lower tropospheric levels. Conversely, when aerosols are prescribed and constant, finer particles are considered and less activation of aerosols to CCN is produced, which lowers the velocity of condensational growth,
375 leading to lower rates of latent heat production and suppression of warm rain provided the difficulty of small aerosols to grow up to droplet size. Hence, the coupling of the meteorological model to an online chemistry module ~~results~~ seems to result of paramount significance for the formation and evolution of a medicane. An interactive calculation of aerosols provides realistic SSA concentrations which, combined with the ability of the model to introduce hygroscopic and microphysical properties
380 for the different species, favours ~~deep convection development~~ the simulation of intense deep convection processes, which are crucial in this type of storms.

Initialization time largely modulates ~~medicane simulations output~~ the output of the medicane simulations, thus being a source of great variability. Highly influenced by the initial conditions, these simulations are prone to lose the necessary conditions for a medicane to be triggered and maintained. Despite this sensitivity, there seems to be no privileged run-up time to simulate
385 medicanes, and no systematic deviation is produced by this factor. Hence, modifying the initialization time is analogous to perturbing the initial conditions in a system highly sensitive to initial conditions. This result ~~is important~~ needs to be highlighted,

as the different run-up times can thus be regarded as an ensemble of perturbed initial conditions, and thus the results obtained related to the importance of an interactive aerosols calculation are robust within this ensemble.

390 Spectral nudging leads to ~~shorter and longer but~~ less intense medicane tracks. ~~Given the deep~~ Besides the fact that it seems useful for producing 'realistic' medicanes in some cases, like that of Celeno, the storms do not seem to be fully developed when SN is introduced. Given the vertical character of medicane structures, any forcing introduced by synoptic scale dynamics may provoke a misalignment in the medicane core or a break in the deep convective structure. Specifically, the asymmetric upper-lower tropospheric forcing derived from the spectral nudging technique without in-PBL influence deviates the vertical alignment of medicane core, introducing an artificial vertical shear that hampers the formation of deep convection. However, 395 further analysis on the wavelength and in-PBL spectral nudging ~~is~~ would be required to completely determine whether spectral nudging could be beneficial for medicane simulations.

Finally, this contribution discusses the differences between simulations with and without SSA feedbacks, but no assessment is provided on whether differences imply a better agreement with actual medicanes. The main reason for this is the lack of reliable observations ~~in-over~~ the sea to carry out a comprehensive validation of the simulations. Therefore the contribution 400 focuses on the physical mechanisms reproduced by the model, and provides arguments ~~that support~~ supporting their feasibility. Given the great similarities of medicanes with tropical cyclones on their mature stage, it seems clear that the possibility for medicanes to produce its own SSA within a feedback process needs to be accounted by the simulations. The analysis included here indeed shows that enabling this possibility leads to deeper and longer medicane tracks. Therefore the natural question emerging is to what extent this deepening of the storm is realistic. Although the lack of observations to carry out validations 405 hampers such an assessment, including these processes can only lead to a more realistic simulation of medicanes. Models, and specifically their microphysics ~~processes~~ or the presence and effects of aerosols, are strongly parameterized. Even if an eventual validation could demonstrate that simulations without SSA feedback, i.e. shallower storms, are closer to observations, this would not mean that ignoring this feedback produces better results. Instead, it would demonstrate that the model is heavily tuned to produce better results when important feedback processes are ignored. Therefore, it becomes evident that the inclusion 410 of SSA feedback is a fundamental mechanism in the development of tropical-like storms, and simulations aiming at studying these phenomena should not neglect its importance.

Author contributions. EPS carried out the simulations and performed the calculations of this paper. JPM contributed to the design of the simulations and their analysis. He also provided ideas for new approaches in the analysis of the simulations that have been integrated in the final manuscript. JJGN, PJG and JPM provided substantial expertise on the topic that contributed to its understanding. The paper has been 415 written by EPS, JJGN and JPM, and all authors have contributed reviewing the text.

Competing interests. The authors declare no conflict of interest.

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References

- WPS V4 Geographical Static Data Downloads Page, https://www2.mmm.ucar.edu/wrf/users/download/get_sources_wps_geog.html, accessed: 2019-12-04.
- NOAA 2011 Tropical Bulletin Archive, <https://www.ssd.noaa.gov/PS/TROP/DATA/2011/bulletins/med/20111108000001M.html>, last accessed: 2020-12-22.
- 425 Akhtar, N., Brauch, J., Dobler, A., Béranger, K., and Ahrens, B.: Medicanes in an ocean–atmosphere coupled regional climate model, *Natural Hazards and Earth System Sciences*, 14, 2189–2201, <https://doi.org/10.5194/nhess-14-2189-2014>, <https://nhess.copernicus.org/articles/14/2189/2014/>, 2014.
- Berrisford, P., Dee, D., Poli, P., Brugge, R., Fielding, M., Fuentes, M., Kållberg, P., Kobayashi, S., Uppala, S., and Simmons, A.: The
430 ERA-Interim archive Version 2.0, p. 23, <https://www.ecmwf.int/node/8174>, 2011.
- Bian, H., Froyd, K., Murphy, D. M., Dibb, J., Darmenov, A., Chin, M., Colarco, P. R., da Silva, A., Kucsera, T. L., Schill, G., et al.:
Observationally constrained analysis of sea salt aerosol in the marine atmosphere, *Atmospheric Chemistry and Physics*, 19, 10773–
10785, 2019.
- Boucher, O., Randall, D., Artaxo, P., Bretherton, C., Feingold, G., Forster, P., Kerminen, V.-M., Kondo, Y., Liao, H., Lohmann, U., et al.:
435 Clouds and aerosols, in: *Climate change 2013: the physical science basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, pp. 571–657, Cambridge University Press, 2013.
- Bouin, M.-N. and Lebeaupin Brossier, C.: Surface processes in the 7 November 2014 medicane from air–sea coupled high-resolution numerical modelling, *Atmospheric Chemistry and Physics*, 20, 6861–6881, 2020.
- Cavicchia, L. and Von Storch, H.: The simulation of Medicanes in a high-resolution regional climate model, *Climate Dynamics*, 39,
440 <https://doi.org/10.1007/s00382-011-1220-0>, 2011.
- Cavicchia, L., von Storch, H., and Gualdi, S.: A long-term climatology of medicanes, *Climate dynamics*, 43, 1183–1195, 2014.
- Cioni, G.: Thermal structure and dynamical modelling of a Mediterranean Tropical-like Cyclone, Ph.D. thesis, <https://doi.org/10.13140/RG.2.1.2000.4722>, 2014.
- Cioni, G., Malguzzi, P., and Buzzi, A.: Thermal structure and dynamical precursor of a Mediterranean tropical-like cyclone, *Quarterly Journal*
445 *of the Royal Meteorological Society*, 142, 1757–1766, 2016.
- Dafis, S., Rysman, J.-F., Claud, C., and Flaounas, E.: Remote sensing of deep convection within a tropical-like cyclone over the Mediterranean
Sea, *Atmospheric Science Letters*, 19, e823, <https://doi.org/10.1002/asl.823>, <https://rmets.onlinelibrary.wiley.com/doi/abs/10.1002/asl.823>, 2018.
- Dafis, S., Claud, C., Kotroni, V., Lagouvardos, K., and Rysman, J.-F.: Insights into the convective evolution of Mediterranean tropical-
450 like cyclones, *Quarterly Journal of the Royal Meteorological Society*, 146, 4147–4169, <https://doi.org/10.1002/qj.3896>, <https://rmets.onlinelibrary.wiley.com/doi/abs/10.1002/qj.3896>, 2020.
- Emanuel, K.: Genesis and maintenance of “Mediterranean hurricanes”, *Advances in Geosciences*, 2, 217–220, <https://hal.archives-ouvertes.fr/hal-00296883>, 2005.
- Fan, J., Wang, Y., Rosenfeld, D., and Liu, X.: Review of aerosol–cloud interactions: Mechanisms, significance, and challenges, *Journal of*
455 *the Atmospheric Sciences*, 73, 4221–4252, 2016.

- Fita, L., Romero, R., Luque, A., Emanuel, K., and Ramis, C.: Analysis of the environments of seven Mediterranean tropical-like storms using an axisymmetric, nonhydrostatic, cloud resolving model, *Natural Hazards and Earth System Science*, 7, 41–56, <https://hal.archives-ouvertes.fr/hal-00301676>, 2007.
- 460 Gaertner, M. Á., González-Alemán, J. J., Romera, R., Domínguez, M., Gil, V., Sánchez, E., Gallardo, C., Miglietta, M. M., Walsh, K. J., Sein, D. V., et al.: Simulation of medicanes over the Mediterranean Sea in a regional climate model ensemble: impact of ocean–atmosphere coupling and increased resolution, *Climate dynamics*, 51, 1041–1057, 2018.
- Gerber, H. E.: Relative-humidity parameterization of the Navy Aerosol Model (NAM), Tech. rep., Naval Research Lab Washington DC, 1985.
- Gong, S., Barrie, L., and Blanchet, J.-P.: Modeling sea-salt aerosols in the atmosphere: 1. Model development, *Journal of Geophysical Research: Atmospheres*, 102, 3805–3818, 1997.
- 465 Gong, S. L.: A parameterization of sea-salt aerosol source function for sub- and super-micron particles, *Global Biogeochemical Cycles*, 17, <https://doi.org/10.1029/2003GB002079>, <https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2003GB002079>, 2003.
- Grell, G. A. and Dévényi, D.: A generalized approach to parameterizing convection combining ensemble and data assimilation techniques, *Geophysical Research Letters*, 29, 38–1, 2002.
- 470 Grell, G. A., Peckham, S. E., Schmitz, R., McKeen, S. A., Frost, G., Skamarock, W. C., and Eder, B.: Fully coupled “online” chemistry within the WRF model, *Atmospheric Environment*, 39, 6957–6975, 2005.
- Hoarau, T., Barthe, C., Tulet, P., Claeys, M., Pinty, J.-P., Bousquet, O., Delanoë, J., and Vié, B.: Impact of the generation and activation of sea salt aerosols on the evolution of Tropical Cyclone Dumile, *Journal of Geophysical Research: Atmospheres*, 123, 8813–8831, 2018.
- Hong, S.-Y., Noh, Y., and Dudhia, J.: A new vertical diffusion package with an explicit treatment of entrainment processes, *Monthly weather review*, 134, 2318–2341, 2006.
- 475 Jensen, J. B. and Nugent, A. D.: Condensational growth of drops formed on giant sea-salt aerosol particles, *Journal of the atmospheric sciences*, 74, 679–697, 2017.
- Jiang, B., Lin, W., Li, F., and Chen, B.: Simulation of the effects of sea-salt aerosols on cloud ice and precipitation of a tropical cyclone, *Atmospheric Science Letters*, 20, e936, 2019a.
- 480 Jiang, B., Lin, W., Li, F., and Chen, J.: Sea-salt aerosol effects on the simulated microphysics and precipitation in a tropical cyclone, *Journal of Meteorological Research*, 33, 115–125, 2019b.
- Jiang, B., Wang, D., Shen, X., Chen, J., and Lin, W.: Effects of sea salt aerosols on precipitation and upper troposphere/lower stratosphere water vapour in tropical cyclone systems, *Scientific reports*, 9, 1–13, 2019c.
- Jiménez, P. A. and Dudhia, J.: Improving the representation of resolved and unresolved topographic effects on surface wind in the WRF model, *Journal of Applied Meteorology and Climatology*, 51, 300–316, 2012.
- 485 Köhler, H.: The nucleus in and the growth of hygroscopic droplets, *Transactions of the Faraday Society*, 32, 1152–1161, 1936.
- Lagouvardos, K., Kotroni, V., Nickovic, S., Jovic, D., Kallos, G., and Tremback, C.: Observations and model simulations of a winter sub-synoptic vortex over the central Mediterranean, *Meteorological Applications: A journal of forecasting, practical applications, training techniques and modelling*, 6, 371–383, 1999.
- 490 Lorente-Plazas, R., Montávez, J. P., Jimenez, P. A., Jerez, S., Gómez-Navarro, J. J., García-Valero, J. A., and Jimenez-Guerrero, P.: Characterization of surface winds over the Iberian Peninsula, *International Journal of Climatology*, 35, 1007–1026, <https://doi.org/https://doi.org/10.1002/joc.4034>, <https://rmets.onlinelibrary.wiley.com/doi/abs/10.1002/joc.4034>, 2015.

- Luo, H., Jiang, B., Li, F., and Lin, W.: Simulation of the effects of sea-salt aerosols on the structure and precipitation of a developed tropical cyclone, *Atmospheric Research*, 217, 120–127, 2019.
- 495 Miglietta, M. M. and Rotunno, R.: Development mechanisms for Mediterranean tropical-like cyclones (medicanes), *Quarterly Journal of the Royal Meteorological Society*, 145, 1444–1460, 2019.
- Miglietta, M. M., Laviola, S., Malvaldi, A., Conte, D., Levizzani, V., and Price, C.: Analysis of tropical-like cyclones over the Mediterranean Sea through a combined modeling and satellite approach, *Geophysical Research Letters*, 40, 2400–2405, <https://doi.org/10.1002/grl.50432>, <https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1002/grl.50432>, 2013.
- 500 Miglietta, M. M., Mastrangelo, D., and Conte, D.: Influence of physics parameterization schemes on the simulation of a tropical-like cyclone in the Mediterranean Sea, *Atmospheric Research*, 153, 360–375, 2015.
- Miguez-Macho, G., Stenchikov, G. L., and Robock, A.: Spectral nudging to eliminate the effects of domain position and geometry in regional climate model simulations, *Journal of Geophysical Research: Atmospheres*, 109, <https://doi.org/10.1029/2003JD004495>, <https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2003JD004495>, 2004.
- 505 Mitchell, K.: The community Noah land-surface model (LSM), User's Guide. Recovered from ftp://ftp.emc.ncep.noaa.gov/mmb/gcp/ldas/noahslm/ver_2, 7, 2005.
- Mlawer, E. J., Taubman, S. J., Brown, P. D., Iacono, M. J., and Clough, S. A.: Radiative transfer for inhomogeneous atmospheres: RRTM, a validated correlated-k model for the longwave, *Journal of Geophysical Research: Atmospheres*, 102, 16 663–16 682, 1997.
- Monin, A. S. and Obukhov, A. M.: Basic laws of turbulent mixing in the surface layer of the atmosphere, *Contrib. Geophys. Inst. Acad. Sci. USSR*, 151, e187, 1954.
- 510 Morrison, H., Thompson, G., and Tatarskii, V.: Impact of cloud microphysics on the development of trailing stratiform precipitation in a simulated squall line: Comparison of one-and two-moment schemes, *Monthly weather review*, 137, 991–1007, 2009.
- Mylonas, M. P., Douvis, K. C., Polychroni, I. D., Politi, N., and Nastos, P. T.: Analysis of a Mediterranean Tropical-Like Cyclone. Sensitivity to WRF Parameterizations and Horizontal Resolution, *Atmosphere*, 10, 425, 2019.
- 515 Noyelle, R., Ulbrich, U., Becker, N., and Meredith, E. P.: Assessing the impact of sea surface temperatures on a simulated medicane using ensemble simulations, *Natural Hazards and Earth System Sciences*, 19, 941–955, 2019.
- Pravia-Sarabia, E., Gómez-Navarro, J. J., Jiménez-Guerrero, P., and Montávez, J. P.: TITAM (v1.0): the Time-Independent Tracking Algorithm for Medicanes, *Geoscientific Model Development*, 13, 6051–6075, <https://doi.org/10.5194/gmd-13-6051-2020>, <https://gmd.copernicus.org/articles/13/6051/2020/>, 2020.
- 520 Pytharoulis, I.: Analysis of a Mediterranean tropical-like cyclone and its sensitivity to the sea surface temperatures, *Atmospheric Research*, 208, 167–179, 2018.
- Pytharoulis, I., Craig, G. C., and Ballard, S. P.: The hurricane-like Mediterranean cyclone of January 1995, *Meteorological Applications*, 7, 261–279, <https://doi.org/10.1017/S1350482700001511>, <https://rmets.onlinelibrary.wiley.com/doi/abs/10.1017/S1350482700001511>, 2000.
- 525 Pytharoulis, I., Kartsios, S., Tegoulas, I., Feidas, H., Miglietta, M. M., Matsangouras, I., and Karacostas, T.: Sensitivity of a mediterranean tropical-like cyclone to physical parameterizations, *Atmosphere*, 9, 436, 2018.
- Ragone, F., Mariotti, M., Parodi, A., Von Hardenberg, J., and Pasquero, C.: A climatological study of western mediterranean medicanes in numerical simulations with explicit and parameterized convection, *Atmosphere*, 9, 397, 2018.

- Reale, O. and Atlas, R.: Tropical Cyclone–Like Vortices in the Extratropics: Observational Evidence and Synoptic Analysis, *Weather and Forecasting*, 16, 7 – 34, [https://doi.org/10.1175/1520-0434\(2001\)016<0007:TCLVIT>2.0.CO;2](https://doi.org/10.1175/1520-0434(2001)016<0007:TCLVIT>2.0.CO;2), https://journals.ametsoc.org/view/journals/wefo/16/1/1520-0434_2001_016_0007_tclvit_2_0_co_2.xml, 01 Feb. 2001.
- Ricchi, A., Miglietta, M. M., Barbariol, F., Benetazzo, A., Bergamasco, A., Bonaldo, D., Cassardo, C., Falcieri, F. M., Modugno, G., Russo, A., et al.: Sensitivity of a Mediterranean tropical-like cyclone to different model configurations and coupling strategies, *Atmosphere*, 8, 92, 2017.
- 535 Rizza, U., Canepa, E., Ricchi, A., Bonaldo, D., Carniel, S., Morichetti, M., Passerini, G., Santiloni, L., Scremin Puhales, F., and Miglietta, M. M.: Influence of wave state and sea spray on the roughness length: Feedback on medicanes, *Atmosphere*, 9, 301, 2018.
- Rizza, U., Canepa, E., Miglietta, M. M., Passerini, G., Morichetti, M., Mancinelli, E., Virgili, S., Besio, G., De Leo, F., and Mazzino, A.: Evaluation of drag coefficients under medicane conditions: Coupling waves, sea spray and surface friction, *Atmospheric Research*, 247, 105 207, <https://doi.org/10.1016/j.atmosres.2020.105207>, <https://www.sciencedirect.com/science/article/pii/S0169809520311431>, 2021.
- 540 Rosenblatt, M.: Remarks on Some Nonparametric Estimates of a Density Function, *Ann. Math. Statist.*, 27, 832–837, <https://doi.org/10.1214/aoms/1177728190>, 1956.
- Rosenfeld, D., Woodley, W. L., Khain, A., Cotton, W. R., Carrió, G., Ginis, I., and Golden, J. H.: Aerosol effects on microstructure and intensity of tropical cyclones, *Bulletin of the American meteorological Society*, 93, 987–1001, 2012.
- Skamarock, W., Klemp, J., Dudhia, J., Gill, D., Barker, D., Wang, W., and Powers, J.: A Description of the Advanced Research WRF Version 3, 27, 3–27, 2008.
- 545 Tous, M. and Romero, R.: Meteorological environments associated with medicane development, *International Journal of Climatology*, 33, 1–14, 2013.
- Tous, M., Romero, R., and Ramis, C.: Surface heat fluxes influence on medicane trajectories and intensification, *Atmospheric research*, 123, 400–411, 2013.