

1 A Comparative Study to Reveal the Influence of Typhoons on the 2 Transport, Production and Accumulation of O₃ in the Pearl River 3 Delta, China

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14 **Abstract.** The Pearl River Delta (PRD) region in South China is faced with severe ambient O₃ pollution in autumn and summer,
15 which mostly coincides with the occurrence of typhoons above the Northwest Pacific. With increasingly severe O₃ pollution
16 in the PRD under the influence of typhoons, it is necessary to gain a comprehensive understanding of the impact of typhoons
17 on O₃ transport, production and accumulation for efficient O₃ reduction. In this study, we analysed the general influence of
18 typhoons on O₃ pollution in the PRD via systematic comparisons of meteorological conditions, O₃ processes and sources on
19 O₃ pollution days with and without typhoon occurrence (denoted as the typhoon-induced and no-typhoon scenarios,
20 respectively), and also examined the differences in these influences in autumn and summer. The results show that the approach
21 of typhoons was accompanied by higher wind speeds and strengthened downdrafts in autumn as well as the inflows of more
22 polluted air masses in summer, suggesting favourable O₃ transport conditions in the typhoon-induced scenario in both seasons.
23 However, the effect of typhoons on the production and accumulation of O₃ were distinct. Typhoons led to reduced cloud cover,
24 and thus stronger solar radiation in autumn, which accelerated O₃ production, but the shorter residence time of local air masses
25 was unfavourable for the accumulation of O₃ within the PRD. In contrast, in summer, typhoons increased cloud cover, and
26 weakened solar radiation, thus restraining O₃ formation, but the growing residence time of local air masses favoured O₃
27 accumulation. The modelling results using the Community Multiscale Air Quality (CMAQ) model for the typical O₃ pollution
28 days suggest increasing contributions from the transport processes as well as sources outside the PRD for O₃ pollution,
29 confirming enhanced O₃ transport under typhoon influence in both seasons. The results of the process analysis in CMAQ
30 suggest that the chemical process contributed more in autumn but less in summer in the PRD. Since O₃ production and
31 accumulation cannot be enhanced at the same time, the proportion of O₃ contributed by emissions within the PRD was likely
32 to decrease in both seasons. The difference in the typhoon influence on O₃ processes in autumn and summer can be attributed
33 to the seasonal variation of the East Asian monsoon. From the “meteorology-process-source” perspective, this study revealed

34 the complex influence of typhoons on O₃ pollution in the PRD and their seasonal differences. To alleviate O₃ pollution under
35 typhoon influence, emission control is needed on a larger scale, rather than only within the PRD.

36 **1 Introduction**

37 Tropospheric ozone (O₃) serves as a secondary pollutant in ambient air and is detrimental for human health and crop
38 production (Wang et al., 2017; Liu et al., 2018; Mills et al., 2018). Ambient O₃ is produced from its precursors, i.e., nitrogen
39 oxides (NO_x = NO + NO₂) and volatile organic compounds (VOCs), through chemical reactions in the presence of sunlight.
40 ~~This O₃ can accumulate locally, or be transported to downwind regions.~~ Due to the relatively long lifetime of O₃ (~22 days;
41 Stevenson et al., 2006), it can accumulate locally, or be transported to downwind regions. Under unfavourable
42 meteorological conditions, enhanced transport, production and/or accumulation of O₃ can all contribute to the O₃ pollution
43 within a region (National Research Council, 1991).

44
45 As the largest city cluster in South China, the Pearl River Delta (PRD) region is faced with frequent ambient O₃ pollution,
46 especially in autumn and summer (Li et al., 2014; Wang et al., 2017; Lu et al., 2018). Along with the continuous increasing of
47 O₃ levels in recent years (Li et al., 2019), O₃ has become the primary contributor to the deterioration of air quality in this
48 region (Feng et al., 2019). The occurrence of O₃ pollution in the PRD is predominantly related to the influence of typhoons
49 (or tropical cyclones) above the Northwest Pacific (Gao et al., 2018; Deng et al., 2019; Lin et al., 2019). According to Gao et
50 al. (2018), seven out of the nine most severe O₃ episodes (regional-mean maximum 8-h average O₃ concentrations ~~of~~ > 240
51 μg/m³) during 2014–2016 coincided with the approach of typhoons. The changes in the track and intensity of typhoons may
52 contribute to the growing trend of O₃ levels recently and in future (Lam, 2018; Lam et al., 2018). Therefore, a
53 comprehensive understanding of the influence of typhoons on the transport, production and accumulation of O₃ has
54 important implications for efficient and strategic O₃ reduction in the PRD.

55
56 Analyses of typhoon-related O₃ episodes in the PRD have been extensively reported in previous publications. The effect of
57 typhoons on O₃ pollution is closely linked to meteorological conditions that are conducive to the transport, production and/or
58 accumulation of O₃. Stagnation caused by typhoons, characterised by low wind speeds, has been reported during many
59 episodes, and it promotes the accumulation of locally formed O₃ within the PRD (Wang et al., 1998; So and Wang, 2003;
60 Wang and Kwok, 2003; Huang et al., 2005; Lam et al., 2005; Jiang et al., 2008; Zhang et al., 2014; Chow et al., 2019).
61 Strong north or west winds were observed or simulated during several episodes, suggesting the potentially strengthened
62 transport of pollutants under typhoon influence (Wang et al., 2001; Yang et al., 2012; Wang et al., 2015; Wei et al., 2016).
63 Downdrafts on the outskirts of typhoons may promote downward O₃ transport and contribute to near-ground O₃ pollution as
64 well (Lam, 2018), but its appearance in the PRD has only been examined in a few studies. Cloudless conditions and strong
65 solar radiation enhance O₃ production, which is another important cause of O₃ pollution (Wang et al., 1998; Wang and

66 Kwok, 2003; Li et al., 2018; Yue et al., 2018; Chow et al., 2019). In a more direct way, several studies have utilised
67 chemical transport models, along with the Process Analysis (PA) tool and source apportionment (SA) methods, to quantify
68 and compare the contributions of various O₃ processes (e.g., transport and the chemical process) and sources (e.g., local
69 emissions, outside emissions and background) during these episodes. Based on reports by Huang et al. (2005), Lam et al.
70 (2005), Jiang et al. (2008), Wang et al. (2010), Li (2013), Wang et al. (2015), Wei et al. (2016) and Chen et al. (2018),
71 horizontal/vertical transport and chemical production may both be the main contributing process for typhoon-induced O₃
72 pollution in different parts of the PRD. The SA results revealed that emissions within the PRD contributed 40–80% of O₃
73 during typhoon-related O₃ episodes (Li et al., 2012; Li, 2013; Chen et al., 2015), suggesting the potentially important role of
74 O₃ accumulation for O₃ pollution here. However, despite massive episode-based studies, several important questions still
75 remain: Are O₃ transport, production and accumulation within the PRD all enhanced at the same time by typhoons? Do both
76 O₃ pollution seasons (autumn and summer) experience similar impact of typhoons on O₃ pollution? More thorough
77 investigations are needed to answer these questions.

78

79 In this study, we present systematic comparisons between O₃ pollution in the typhoon-induced and no-typhoon scenarios
80 (definitions given in Sect. 2.2) to elucidate the influence of typhoons on O₃ transport, production and accumulation in the PRD
81 and to reveal their seasonal differences. October and July in 2014–2018 were selected as the representative months for autumn
82 and summer, respectively. Multiple datasets, including the ERA-Interim re-analysis, the routine monitoring datasets,
83 trajectories calculated by the Hysplit model and the modelling results of typical O₃ pollution days using the Community
84 Multiscale Air Quality (CMAQ) model, were used in the comparisons. A detailed introduction of these datasets is presented
85 in Sect. 2. The comparisons were conducted from the perspectives of meteorological conditions (Sect. 3), O₃ processes and
86 sources (Sect. 4), and the conclusions about the influence of typhoons on the causes of ambient O₃ pollution in the PRD in the
87 two seasons are illustrated in Sect. 5.

88 **2 Methods**

89 **2.1 Datasets**

90 The detailed information for the datasets utilised in the comparison of meteorological conditions is presented below:

- 91 • **Re-analysis datasets:** We mainly used the ERA-Interim re-analysis product in the analyses due to its more available
92 parameters and high spatial coverage (available at [https://www.ecmwf.int/en/forecasts/datasets/reanalysis-datasets/era-](https://www.ecmwf.int/en/forecasts/datasets/reanalysis-datasets/era-interim)
93 [interim](https://www.ecmwf.int/en/forecasts/datasets/reanalysis-datasets/era-interim), last accessed: March 2020; Dee et al., 2011; Berrisford et al., 2011). Specifically, meteorological parameters
94 used in the comparisons include the following three categories: (1) near-surface parameters from the analysis fields,
95 including air temperature (at a height of 2 m), relative humidity (RH, at 1000 hPa), horizontal wind speeds (at a height
96 of 10 m; zonal and meridional wind speeds were also involved in the comparisons), and low (for the height at which
97 pressure/surface pressure > 0.8), medium (for the height at which 0.45 < pressure/surface pressure < 0.8), high (for the

98 height at which pressure/surface pressure < 0.45) and total cloud covers; (2) near-surface parameters from the forecast
99 fields, including plenary boundary layer (PBL) height and net surface solar radiation; and (3) upper air parameters at
100 multiple heights, including horizontal and vertical wind speeds, cloud water content and O₃ mixing ratio. The focus of
101 this study is O₃ pollution during the daytime, and therefore, only the parameters at 14:00 local time (LT) were selected
102 for the analyses (except for net surface solar radiation, which was averaged within 8:00–17:00 LT).

- 103 • **Surface meteorological routine monitoring datasets:** The routine monitoring meteorological data collected at 29
104 national meteorological sites within the PRD (locations shown in Fig. S1a) were also used to explore the
105 meteorological features under the impact of typhoons. The parameters include air temperature, RH, and wind speed and
106 direction (also transformed to zonal and meridional wind speeds in the comparisons) at 14:00 LT.
- 107 • **Typhoon information:** The times, locations and intensities of typhoons were provided by the Chinese Meteorological
108 Administration Best Track Dataset of tropical cyclones (Ying et al., 2014). The tracks of all typhoons that potentially
109 contributed to O₃ pollution in the PRD during the study period (October and July in 2014–2018) are shown in Fig. S2
110 and S3.
- 111 • **O₃ concentrations:** Hourly O₃ concentration data, which were originally released by the China National Environmental
112 Monitoring Centre, were downloaded from <http://beijingair.sinaapp.com> (last accessed: Dec. 2018). Based on the
113 hourly data, we calculated the maximum 1-hr concentrations (MDA1) and maximum 8-hr average concentrations
114 (MDA8) of O₃ in nine municipalities in the PRD (including Guangzhou, Shenzhen, Zhuhai, Foshan, Jiangmen,
115 Zhaoqing, Huizhou, Dongguan and Zhongshan) to identify O₃ pollution days that served as samples in the comparisons.

116 2.2 Definition and classification of O₃ pollution days

117 In this study, O₃ pollution days were defined as the days when the MDA1 exceeds 200 µg/m³ or the MDA8 exceeds 160
118 µg/m³ for O₃ (both are the Grade-II thresholds of the Chinese National Ambient Air Quality Standard (NAAQS), GB 3095-
119 2012) in any of the nine municipalities in the PRD. According to these criteria, there were 78 and 55 O₃ pollution days
120 (given in Table S1 and S2) during October and July in 2014–2018, respectively. The information about these O₃ pollution
121 days in the two representative months is listed in Table 1 (overall) and S3 (monthly), including the numbers of days, their
122 proportions in the month, and the corresponding mean O₃ concentrations (MDA8 and MDA1, highest values among nine
123 municipalities in the PRD). Although there were more O₃ pollution days in October than in July, ~~O₃ pollution days under~~
124 ~~typhoon influence accounted for ~30% of all days in both months. O₃ pollution under typhoon influence occurred on ~30%~~
125 ~~days of both months.~~ Higher O₃ MDA1 and MDA8 values can be ~~generally~~ found with the appearance of typhoons in
126 comparison with days without typhoons ~~in July, whereas these values are similar in October, further~~ indicating the important
127 role of typhoons in O₃ pollution in the PRD.

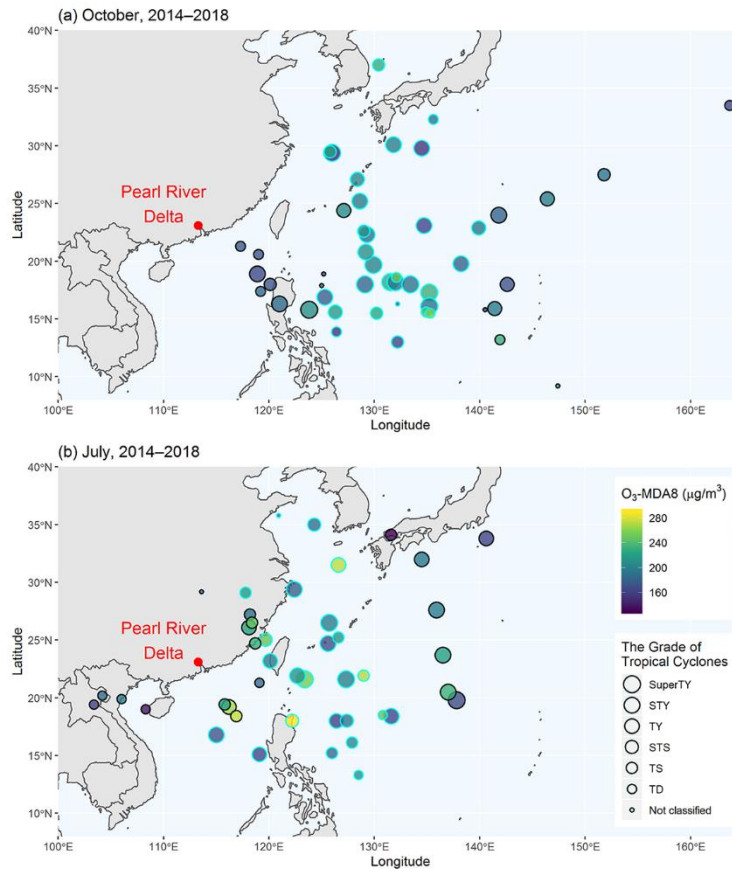
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129 The differing locations of typhoons can result in the diverse effect of typhoons on O₃ pollution (Chow et al., 2018). To
130 determine the general influence of typhoons on O₃ pollution in the PRD, it was necessary to further select O₃ pollution days

131 coinciding with typhoons with similar directions and distances ~~from to~~ the PRD. ~~First, we removed five O₃ pollution days in~~
 132 ~~July with typhoons located to the due north or southwest of the PRD from the analyses.~~ As is shown in Fig. 1, ~~the remaining~~
 133 ~~days, including~~ all O₃ pollution days in October and most O₃ pollution days in July under typhoon influence, were associated
 134 with typhoons to the east of the PRD, which were more likely to cause O₃ pollution (Chow et al., 2018). In order to minimize
 135 the disturbance of typhoon directions in the comparisons, we removed the remaining five O₃ pollution days in July with
 136 typhoons located to the due north or southwest of the PRD from the analyses. After this, based on the distances between
 137 typhoon centres and the PRD (at 14:00 LT), we classified the pollution days with typhoons in each season into three
 138 categories: close typhoon (lowest 20% of distances), typhoon (20–80% intervals of distances), and far typhoon (longest 20%
 139 of distances)-induced days. The typhoon-induced days represent O₃ pollution days with general typhoon influence, and they
 140 were compared with those without the appearance of typhoons (hereafter denoted as the no-typhoon days). It should be noted
 141 that the distances between typhoon centres and the PRD on the typhoon-induced days were overall larger in autumn (1400–
 142 2800 km, at 14:00 LT) than in summer (700–2000 km, at 14:00 LT), which may be the consequence of the different
 143 characteristics of typhoon paths in the two seasons: most typhoons in autumn travel northwest initially and then turn
 144 northward in the areas east of the Philippines (Fig. S2), whereas they are more likely to end up landing in Southeast China in
 145 summer (Fig. S3). Since the influence of typhoons on O₃ pollution may be different when typhoons come close enough to
 146 the PRD (Lam et al., 2005; Li, 2013), the close typhoon-induced days were considered to be a special scenario in the
 147 comparisons of meteorological conditions (Sect. 3.5). Owing to the less apparent effect of typhoons over the PRD, we did
 148 not include the far typhoon-induced days in the discussions.

149 **Table 1.** The numbers and proportions of O₃ pollution days, and O₃ concentrations for various scenarios.

Parameter	October, 2014–2018	July, 2014–2018	150
Number (proportion) of O ₃ pollution days	78 (50.3%)	55 (35.5%)	151
With typhoons	49 (31.6%)	45 (29.0%)	
Typhoon-induced days	30 (19.4%)	24 (15.5%)	
Close typhoon-induced days	10 (6.5%)	8 (5.2%)	
Without typhoons (no-typhoon days)	29 (18.7%)	10 (6.5%)	
Mean PRD-max O ₃ MDA8 (µg/m ³)			
With typhoons	195.0	205.3	
Typhoon-induced days	199.5	205.4	
Close typhoon-induced days	184.6	225.7	
Without typhoons (no-typhoon days)	189.8	187.8	
Mean PRD-max O ₃ MDA1 (µg/m ³)			
With typhoons	230.4	259.8	
Typhoon-induced days	235.2	260.0	
Close typhoon-induced days	219.2	277.1	
Without typhoons (no-typhoon days)	231.5	246.5	



152

153 **Figure 1.** The location and intensity of typhoons at 14:00 LT on all O₃ pollution days with typhoons, and the corresponding O₃ MDA8
 154 concentrations (maximum values in the nine municipalities of the PRD) on the same days during (a) October and (b) July in 2014–2018.
 155 The points with cyan borders indicate the “typhoon-induced” O₃ pollution days used in the comparisons. The grades of tropical cyclones
 156 (Chinese National Standard, GB/T 19201-2006) are as follows: SuperTY - super typhoon; STY - severe typhoon; TY - typhoon; STS -
 157 severe tropical storm; TS - tropical storm; TD - tropical depression; others are grouped as “not classified”.

158 2.3 Calculation of the trajectories and air parcel residence time

159 To explore the potential effect of cross-regional transport on O₃ pollution in the PRD, we applied the Hysplit model (Stein et
 160 al., 2015) with the Global Data Assimilation System (GDAS) datasets as inputs to calculate 72-h backward trajectories reaching
 161 the PRD at 14:00 LT for all O₃ pollution days. The Modiesha site (23.1°N, 113.3°E; Fig. S1b), which is located in the central
 162 part of the PRD, was the endpoint of backward trajectories, with its height set as 500 m above the ground. Its height was set
 163 as 500 m above the ground to better represent the effect of long-range transport on O₃ pollution, as well as to minimize the
 164 disturbance of objects near the surface to the transport (Park et al., 2007).

165

166 Air parcel residence time (APRT), discussed by Huang et al. (2019), is the average number of hours that air parcels originated
 167 from one place stay within a pre-defined domain, and long APRTs can be used to indicate good accumulation conditions for

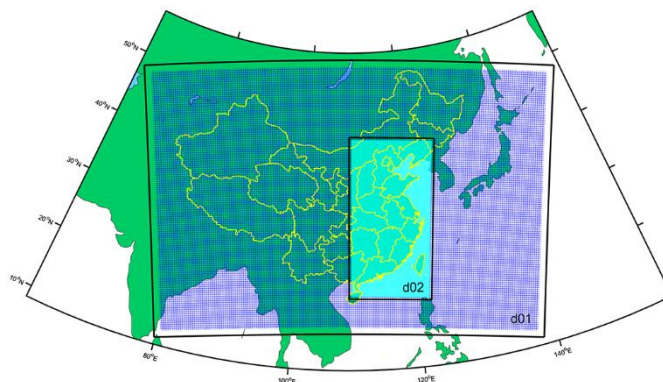
168 locally sourced pollutants. To calculate APRTs in the PRD, we designed a 21×15 point matrix (resolution: 0.2°×0.2°) that
169 embraces the whole PRD (Fig. S4), and forward trajectories starting from these points were also calculated using the Hysplit
170 model. The height of all points was set as 100 m above the ground, ~~which is close to the height of emissions to represent the~~
171 height of all local emissions and to reduce the disturbance of the surface, as well. The start times were set as 2:00, 8:00, 14:00
172 and 20:00 LT for all O₃ pollution days. Afterwards, the length of time each trajectory remained within the administration
173 borders of the PRD, i.e., APRT, was calculated and attributed to its starting point. ~~The APRT values of all points were averaged~~
174 for each scenario and were interpolated to obtain field results. APRTs in each point were averaged, and these averaged APRT
175 values in all points were interpolated using the Kriging method to obtain field results for the further comparisons.

176 2.4 CMAQ modelling: basic setups and modelling methods

177 We utilised the widely used 3D chemical transport model, ~~the CMAQ model~~ (version 5.0.2), to investigate the effects of
178 typhoons on O₃ processes and sources. October 2015 and July 2016 featured the most severe O₃ pollution under typhoon
179 influence among all representative months of the two seasons (Table S3), and thus, they were chosen as the period in the
180 ~~CMAQ~~ modelling (because there was no severe O₃ pollution during the first 10 days of October 2015 and 3–5 November can
181 be classified as the no-typhoon O₃ pollution days, we adjusted the modelling period in autumn to 11 October–10 November
182 2015) and all typhoon-induced and no-typhoon O₃ pollution days in these two months served as representative O₃ pollution
183 days ~~under multiple scenarios in the comparisons~~. In detail, there were four typhoon-induced O₃ pollution days (14–16 and 21
184 October 2015) and four no-typhoon O₃ pollution days (28 October and 3–5 November 2015) in October 2015, whereas there
185 were four and six typhoon-induced and no-typhoon days in July 2016, respectively (typhoon-induced: 7–8 and 30–31 July
186 2016; and no-typhoon: 22–26 and 29 July 2016). The results of ~~the~~ daytime (9:00–17:00 LT) O₃ PA and SA on the above O₃
187 pollution days were averaged for each the typhoon-induced and no-typhoon scenarios in autumn (October 2015) and summer
188 (July 2016) and were used in the comparisons.

189
190 The main setups of the CMAQ modelling are presented as follows. Two-nested modelling domains with the resolutions of 36
191 and 12 km (denoted as d01 and d02, respectively) were set in this study (Fig. 2). Specifically, d02 covers the whole East and
192 Central China (EC-China), enabling us to evaluate the contribution of emissions in these areas to O₃ pollution in the PRD.
193 There were 19 vertical layers in the CMAQ modelling, with about 10 layers within the PBL (about 0–1 km in heights; Guo et
194 al., 2016). The Weather Research and Forecasting (WRF) model (version 3.2) provided the meteorological fields used as inputs.
195 SMOKE (version 2.5) and MEGAN (version 2.10) were used to process the anthropogenic and biogenic emission files,
196 respectively. The anthropogenic emission inventory used in this study consisted of the following three parts: (1) emissions in
197 the PRD, which were provided by the Guangdong Environmental Monitoring Centre; (2) emissions in other areas of mainland
198 China, which were extracted from the MEIC inventory (He, 2012); and (3) emissions in other countries and regions in Asia,
199 which were extracted from the MIX inventory (Li et al., 2017). The initial and boundary conditions of the d01 modelling were
200 obtained from the same-period results of the MOZART-4 global model (available at <https://www.acom.ucar.edu/wrf->

201 chem/mozart.shtml, last accessed: Dec. 2019), and those of the d02 modelling were extracted from the d01 modelling results.
202 The SAPRC07 gas-phase chemistry mechanism (Carter, 2010) and the AERO6 aerosol scheme were set in the **CMAQ**
203 modelling. In addition, the simulations of the two months were both started 10 days ahead to minimise the disturbance of the
204 bias of the initial conditions. The modelling performances of CMAQ and WRF were determined to be acceptable based on the
205 comparisons between the observational and modelling series of meteorological parameters, O₃ MDA8, daily NO₂
206 concentrations and the mixing ratios of non-methane hydrocarbons (NMHCs) in the PRD (for details, refer to Sect. 1 ~~of~~in the
207 Supplement Information), which ensures the validity of the further analyses.
208



209
210 **Figure 2.** Two-nested modelling domain, noted as d01 and d02. The black boxes indicate the WRF modelling domains, and the nested
211 areas are the CMAQ modelling domains.

212 The PA tool in CMAQ was implemented to quantify the hourly contributions of O₃ processes (or integrated process rate, IPR),
213 which includes vertical/horizontal transport (convection+diffusion), chemical process (net O₃ production through gas-phase
214 reactions), dry deposition and cloud process. To explore the overall effect of typhoons on O₃ transport and production in the
215 region, the mean PA results within the administration boundaries of the PRD were calculated and compared.

216
217 In order to identify the sources of all O₃ in the PRD, ~~W~~we used the classic Brute Force Method (BFM) to identify the
218 contributions of emissions (including anthropogenic and biogenic emissions) in the PRD and other regions in the d02 (mainly
219 EC-China), as well as regions outside the d02 (the boundary conditions of the d02) for O₃ pollution in the PRD (hereafter
220 denoted as the contributions of PRD, EC-China and BCON, or S_{PRD} , $S_{EC-China}$ and S_{BCON} , respectively). For a pollutant, the
221 contribution of a specific emission, E_i , can be calculated in two ways: (1) the difference between the modelled concentrations
222 of the base case (all emissions involved) and the sensitivity case where E_i is zeroed out (i.e., top-down BFM); (2) the difference
223 between two sensitivity cases where emissions expect E_i and all of the emissions are zeroed out, respectively (i.e., bottom-up
224 BFM). Owing to the non-linearity between O₃ and its precursors, biases may occur between the results of two types of BFM
225 methods, leading to the non-additivity of the results (Clappier et al., 2017). Therefore, the average of the top-down and bottom-

226 up BFM results was treated as the quantified contributions of the concerned sources. Four simulation cases were run in this
227 study, including (the modelled O₃ concentration in each case was also marked in brackets):

- 228 • the base case (C_{base});
- 229 • the PRD-cut case (C_{PRD_cut}), where emissions within the PRD were zeroed out;
- 230 • the PRD-only case (C_{PRD_only}), where emissions outside the PRD (within d02) were zeroed out; and
- 231 • the zero-emission case (C_0), where all emissions within the d02 were zeroed out.

232 Afterwards, the S_{PRD} , $S_{EC-China}$ and S_{BCON} values (in concentrations) in the polluted areas of the PRD (where modelled daytime
233 O₃ concentrations > 160 µg/m³, the Grade-II O₃ MDA8 thresholds of the Chinese NAAQS) were calculated using the following
234 equations,

$$S_{PRD} = \frac{1}{2} [(C_{base} - C_{PRD_cut}) + (C_{PRD_only} - C_0)], \quad (R1)$$

$$S_{EC-China} = \frac{1}{2} [(C_{base} - C_{PRD_only}) + (C_{PRD_cut} - C_0)], \quad (R2)$$

$$S_{BCON} = C_0. \quad (R3)$$

235 The percentage forms of these values were used in the comparisons.

236 3 Comparison of meteorological conditions

237 3.1 Overview: comparison of meteorological parameters in the PRD

238 First, we compared near-ground meteorological parameters in the PRD on the typhoon-induced and no-typhoon O₃ pollution
239 days. The parameters from ~~the ERA-Interim re-analysis (including the parameters of the first and second categories in Sect.~~
240 ~~2.1) and the~~ routine monitoring datasets (including air temperature, RH, wind speed, zonal and meridional wind speeds
241 measured at 14:00 LT of all O₃ pollution days at 29 national meteorological sites within the PRD (Fig. S1a) and the ERA-
242 Interim re-analysis (including all near-surface parameters from the analysis and forecast fields introduced in Sect. 2.1,
243 extracted at the same time and the locations of sites as these in routine monitoring datasets) were used in the comparison
244 (since all O₃ pollution days in October and over 60% of O₃ pollution days in July were characterized with sunny, cloudy, or
245 overcast weathers with no rainfall in the PRD (Table S4, represented by the weather in Guangzhou), precipitation was not
246 considered in the comparisons). The Mann-Whitney U test was applied to determine whether the above parameters were
247 significantly different ($p < 0.05$) ~~in the two types of O₃ pollution scenarios~~ between typhoon-induced and no-typhoon O₃
248 pollution days.

249

250 As is listed in Table 2, statistically significant differences between the typhoon-induced and no-typhoon scenarios existed for
251 most of the parameters, such as meridional (south-north) wind speed, cloud covers within various height ranges and net
252 surface solar radiation — in both seasons, these parameters were significantly different for the two scenarios. It indicates that
253 the causes of O₃ pollution may vary on typhoon-induced and no-typhoon O₃ pollution days. Note that air temperature, one of

254 the parameters most closely related to O₃ pollution in the PRD (Zhao et al., 2019), was not significantly different in the two
255 scenarios. We also found that the comparison in autumn and summer did not produce the same results: the typhoon-induced
256 days in autumn featured lower RH, stronger winds (especially north wind), reduced cloud cover (low, medium, high and
257 total) and stronger surface solar radiation, whereas in summer, these days had higher RH, weaker south winds, more cloud
258 cover (medium, high and total), weaker surface solar radiation and lower PBL heights. Therefore, the impact of typhoons on
259 O₃ pollution differs in the two seasons, as well. In order to reveal the impact of typhoons on O₃ transport, production, and
260 accumulation in the PRD, more detailed comparisons of the corresponding meteorological indicators are presented in the
261 following sections.

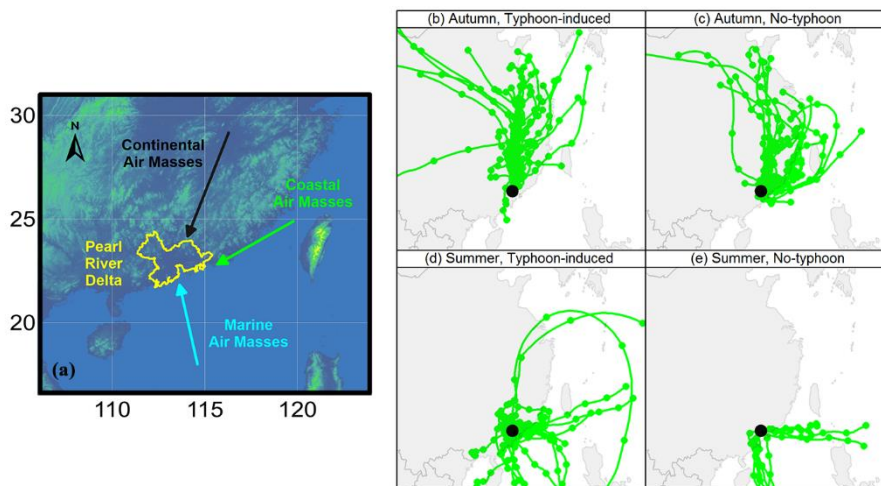
262 **Table 2.** The comparisons of meteorological parameters (all at 14:00 LT except for net surface solar radiation, which is the average value
 263 for 9:00–17:00 LT) in the PRD for the three scenarios (no-typhoon, typhoon-induced, close typhoon-induced) in two seasons (autumn,
 264 summer). RM, routine measurement; ERA, ERA-Interim re-analysis. All of the parameters are presented as “the mean value \pm standard
 265 deviation”. The differences between parameters in the typhoon-induced or close typhoon-induced scenarios and the corresponding
 266 typhoon-induced scenarios for the same season are given in parentheses, and “*” indicates $p < 0.05$, or statistically significant differences
 267 between these parameters when the Mann-Whitney U test is used.

Parameters	Data Source	Autumn (October, 2014–2018)			Summer (July, 2014–2018)		
		No-typhoon	Typhoon-induced	Close Typhoon-induced	No-typhoon	Typhoon-induced	Close Typhoon-induced
Air Temperature (°C)	RM	29.1 \pm 2.2	29.3 \pm 1.8 (+0.2)	29.6 \pm 1.5 (+0.5, *)	33.7 \pm 2.0	33.9 \pm 2.0 (+0.2)	35.0 \pm 1.5 (+1.3, *)
	ERA	29.2 \pm 2.1	29.3 \pm 1.6 (+0.1)	29.6 \pm 1.5 (+0.4, *)	33.4 \pm 1.8	33.5 \pm 1.4 (+0.1)	34.6 \pm 1.4 (+1.2, *)
RH (%)	RM	52.4 \pm 10.2	44.8 \pm 10.4 (-7.6, *)	51.4 \pm 12.4 (-1.0)	57.0 \pm 9.3	58.3 \pm 9.7 (+1.3)	56.9 \pm 6.4 (-0.1)
	ERA	54.0 \pm 9.8	48.3 \pm 11.2 (-5.7, *)	52.2 \pm 12.4 (-1.8, *)	62.6 \pm 10.8	66.4 \pm 9.4 (+3.8, *)	62.5 \pm 9.4 (+0.1)
Wind Speed (m/s)	RM	2.33 \pm 1.18	2.58 \pm 1.23 (+0.25, *)	2.96 \pm 1.40 (+0.63, *)	2.46 \pm 1.33	2.30 \pm 1.20 (-0.16)	2.53 \pm 1.16 (+0.07)
	ERA	2.39 \pm 1.30	2.54 \pm 0.99 (+0.15, *)	3.53 \pm 1.11 (+1.14, *)	2.41 \pm 0.99	2.18 \pm 1.18 (-0.23, *)	2.61 \pm 1.05 (+0.20)
Zonal (East-West) Wind Speed (m/s)	RM	-0.83 \pm 1.72	-0.59 \pm 1.70 (+0.24, *)	-0.13 \pm 1.74 (+0.70, *)	-0.41 \pm 2.05	-0.03 \pm 1.94 (+0.38)	0.73 \pm 1.98 (+1.14, *)
	ERA	-1.41 \pm 1.43	-1.07 \pm 1.04 (+0.34, *)	-0.87 \pm 0.79 (+0.54, *)	0.22 \pm 1.73	-0.02 \pm 1.81 (-0.24)	0.29 \pm 2.45 (+0.07)
Meridional (South-North) Wind Speed (m/s)	RM	-0.36 \pm 1.74	-1.49 \pm 1.66 (-1.13, *)	-2.21 \pm 1.66 (-1.85, *)	0.79 \pm 1.69	0.01 \pm 1.72 (-0.78, *)	-0.69 \pm 1.68 (-1.48, *)
	ERA	-0.27 \pm 1.82	-1.97 \pm 1.16 (-1.70, *)	-3.27 \pm 1.29 (-3.00, *)	1.61 \pm 1.09	0.64 \pm 1.58 (-0.97, *)	-0.68 \pm 1.19 (-2.29, *)
Low Cloud Cover (%)	ERA	17.2 \pm 22.7	4.2 \pm 11.9 (-13.0, *)	15.5 \pm 23.9 (-1.7, *)	8.7 \pm 9.4	7.1 \pm 9.5 (-1.6, *)	5.2 \pm 5.0 (-3.5, *)
Medium Cloud Cover (%)	ERA	22.2 \pm 26.5	10.4 \pm 19.7 (-11.8, *)	9.5 \pm 14.5 (-12.7, *)	8.7 \pm 11.1	15.4 \pm 15.1 (+6.7, *)	21.5 \pm 15.5 (+12.8, *)
High Cloud Cover (%)	ERA	12.1 \pm 23.1	7.2 \pm 16.3 (-4.9, *)	34.6 \pm 35.6 (+22.5, *)	32.2 \pm 30.0	44.9 \pm 29.3 (+12.7, *)	51.0 \pm 34.2 (+18.8, *)
Total Cloud Cover (%)	ERA	43.5 \pm 32.3	20.5 \pm 25.7 (-23.0, *)	51.9 \pm 33.1 (+8.4, *)	43.7 \pm 26.7	58.3 \pm 22.7 (+14.6, *)	67.5 \pm 21.0 (+23.7, *)
Net Surface Solar Radiation (W/m ²)	ERA	456.9 \pm 78.4	516.6 \pm 66.7 (+59.7, *)	516.5 \pm 62.8 (+59.6, *)	560.3 \pm 93.1	523.2 \pm 74.4 (-37.1, *)	541.9 \pm 54.0 (-18.4, *)
PBL Height (m)	ERA	1471 \pm 315	1473 \pm 348 (+2)	1349 \pm 227 (-122, *)	1268 \pm 383	1037 \pm 289 (-231, *)	1196 \pm 300 (-72, *)

268 3.2 O₃ transport conditions: comparison of wind speeds, backward trajectories and vertical air motions

269 The higher wind speeds and/or O₃ levels in the transported air masses are, the more likely O₃ transport plays an increasingly
 270 important role in O₃ pollution. In the PRD, O₃ levels are closely linked to the type of air masses influencing the region, which
 271 can be identified based on backward trajectories. According to Zheng et al. (2010), there are generally three types of air masses

272 that are transported into the PRD along different paths and contribute to O₃ pollution here, namely, the continental, coastal and
 273 marine air masses (Fig. 3a). The continental and coastal air masses can bring O₃ from EC-China to the PRD, and thus, they
 274 are typically recognised as being polluted and contributing to relatively high O₃ levels in the PRD. In contrast, the marine air
 275 masses, originated from the South China Sea, are much cleaner. In this section, we studied the influence of typhoons on O₃
 276 transport by comparing wind speeds and 72-h backward trajectories in ~~various~~ the typhoon-induced and no-typhoon scenarios.
 277



278

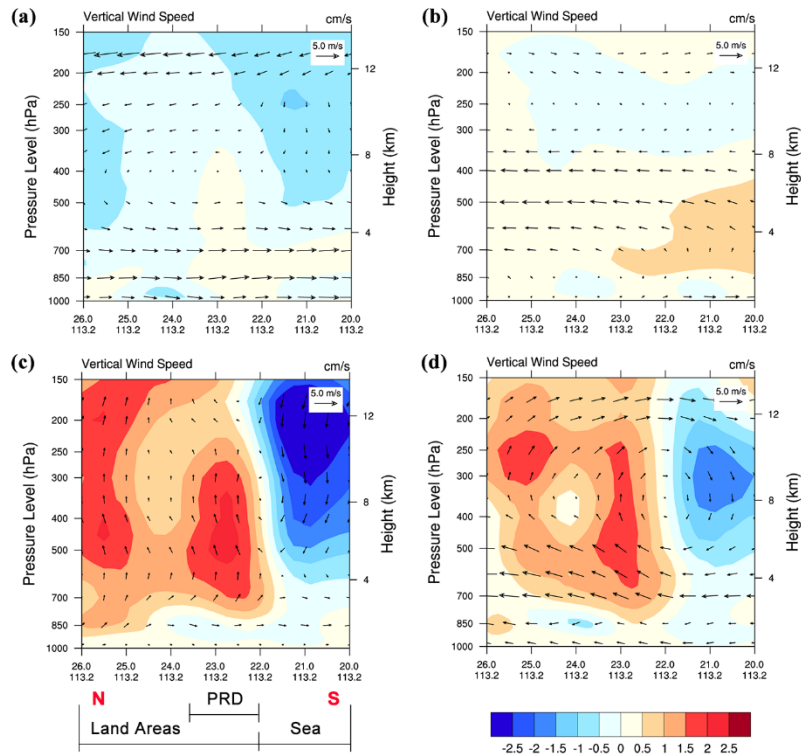
279 **Figure 3.** (a) Three O₃ transport paths towards the PRD. (b–e) Backward trajectories at 14:00 LT for ~~the four scenarios~~ each scenario: (b)
 280 autumn, typhoon-induced; (c) autumn, no-typhoon; (d) summer, typhoon-induced; and (e) summer, no-typhoon. The black dots indicate
 281 the end point of all trajectories, i.e., where the Modiesha site in the central PRD is located.

282 As is displayed in Fig. 3b–c, we identified the influence of continental air masses on the typhoon-induced O₃ pollution days
 283 in autumn, as well as mixed contributions from the continental and coastal air masses on the corresponding no-typhoon days.
 284 However, for the former scenario, significantly increased wind speeds (Table 2) ensure more favourable conditions for the
 285 transport of O₃. In summer, the three types of air masses may all have affected O₃ pollution in the typhoon-induced scenario,
 286 while only the marine air masses influenced the PRD in the no-typhoon scenario (Fig. 3d–e). ~~Since wind speeds did not vary~~
 287 ~~significantly (Table 2), the inflows of much more polluted air masses resulted in that typhoons also tended to increase the~~
 288 ~~contribution of transport to O₃ pollution in the PRD in summer. The increasing influence of much more polluted air masses~~
 289 ~~(continental and coastal air masses) led by typhoon ensured that more O₃ was transported to the PRD, thus typhoons also~~
 290 ~~tended to increase the contribution of transport to O₃ pollution in the PRD in summer.~~ In addition, the influence of different
 291 air masses was also accompanied with variations in the prevailing winds in the PRD, that is, north winds and easterlies in the
 292 typhoon-induced and no-typhoon scenarios in autumn, respectively, and southwest winds in the no-typhoon scenario in
 293 summer (indicated by wind roses in Fig. S5). For the typhoon-induced scenario in summer, the dominate wind direction is
 294 hard to determine. These variations in the local wind fields potentially result in the different spatial distribution of O₃
 295 concentrations in various scenarios.

297 Downdrafts are typically considered to be an important cause of typhoon-induced O₃ pollution (Lam, 2018), but in which
298 scenarios downdrafts influence the PRD remains unclear. Thus, we explored the overall features of vertical air motions from
299 the surface layer to the tropopause in ~~four~~ the typhoon-induced and no-typhoon scenarios, and the ERA-Interim reanalysis
300 dataset (including ~~the parameters of the third category~~ all upper air parameters at multiple heights introduced in Sect. 2.1) was
301 utilised in the comparisons. The contours in Fig. 4 show the cross sections of mean vertical wind speed at 14:00 LT of all O₃
302 pollution days corresponding to the typhoon-induced and no-typhoon scenarios of two seasons, which were made along the
303 113.2°E longitude line, from 26.0°N to 20.0°N ~~along the 113.2°E longitude line~~ (Fig. S4). On the typhoon-induced days in
304 autumn, downdrafts occurred over large areas above the PRD, especially above a height of ~~4 km~~ 700 hPa. Although updrafts
305 can still be found near the sea surface in this scenario, vertical wind speeds tended to be lower compared with those on the no-
306 typhoon days in autumn, which also suggests the enhancement of downdrafts caused by typhoons. In summer, the influence
307 of downdrafts was found over the PRD under 850 hPa on the typhoon-induced O₃ pollution days. However, overall, updrafts
308 prevailed above the land areas and downdrafts prevailed above the sea in both the typhoon-induced and no-typhoon scenarios
309 in summer, which is recognised as the structure of the East Asian summer monsoon cell (Chen et al., 1964; Jin et al., 2013;
310 Ding et al., 2018). For both updrafts and downdrafts, the absolute values of vertical wind speeds in the typhoon-induced
311 scenario in summer were overall higher than these in the corresponding no-typhoon scenario. Therefore, the approach of
312 typhoons did not break the structure of the summer monsoon cell, but rather they further strengthened the vertical motions
313 above both land areas and sea. These analyses suggest that typhoons do not necessarily lead to downdrafts during O₃ pollution
314 periods in the PRD and its adjacent areas; and in summer, vertical air motions affected by typhoons are more complicated than
315 expected owing to the existence of the East Asian summer monsoon.

316

317 We also explored the regions where downdrafts and updrafts occurred on a larger scale and their potential connections with
318 O₃ levels. As is shown in Fig. 5, though updrafts appeared in the PRD at 850 hPa on the typhoon-induced days in autumn,
319 downdrafts dominated in the region at 700 and 500 hPa. For the areas to the north of the PRD, the important role of downdrafts
320 was found at all three heights. In contrast to the no-typhoon days in autumn, downdrafts tended to cover much larger areas in
321 this scenario. Moreover, these areas at 850 and 700 hPa generally featured higher O₃ mixing ratios as well as lower RH (Fig.
322 S6) than others, which is a sign of possible direct downward O₃ transport (Roux et al., 2020; Wang et al., 2020). This part of
323 O₃ can notably aggravate near-ground O₃ pollution in the PRD. In contrast, in summer, updrafts dominated the PRD at various
324 heights in both scenarios. Besides the PRD, most of the regions near the coast were characterised by updrafts above the land
325 as well as downdrafts offshore, further indicating the ubiquity of the summer monsoon cell. By comparing the two scenarios
326 in summer, we found that typhoons resulted in more areas being influenced by updrafts. The areas with high O₃ levels did not
327 coincide with the downdraft-affected areas, and therefore, O₃ transported from the upper air may play a less significant role in
328 the typhoon-induced O₃ pollution in summer.



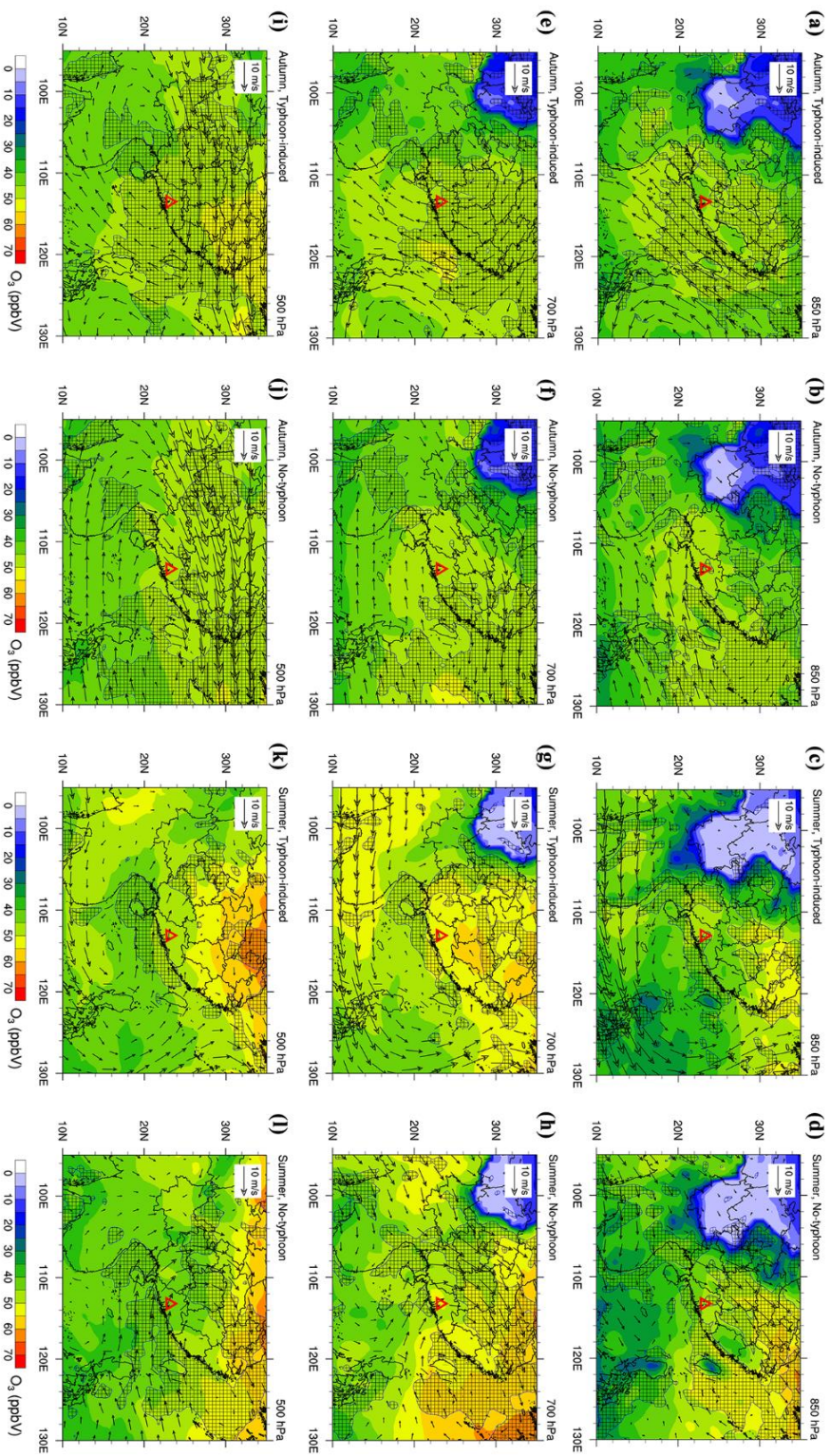
329

330 **Figure 4.** The cross sections of mean vertical wind field at 14:00 LT for ~~the four each~~ scenarios: (a) autumn, typhoon-induced; (b) autumn,
 331 no-typhoon; (c) summer, typhoon-induced; and (d) summer, no-typhoon. Cross sections ~~are were made along the 113.2 E longitude line,~~
 332 ~~from 26.0 N to 20.0 N along the 113.2 E longitude line~~ (Fig. S4). The vectors indicate meridional wind speed (m/s) and vertical wind
 333 speed (cm/s), and the contours indicate vertical wind speed (cm/s). PRD, the Pearl River Delta.

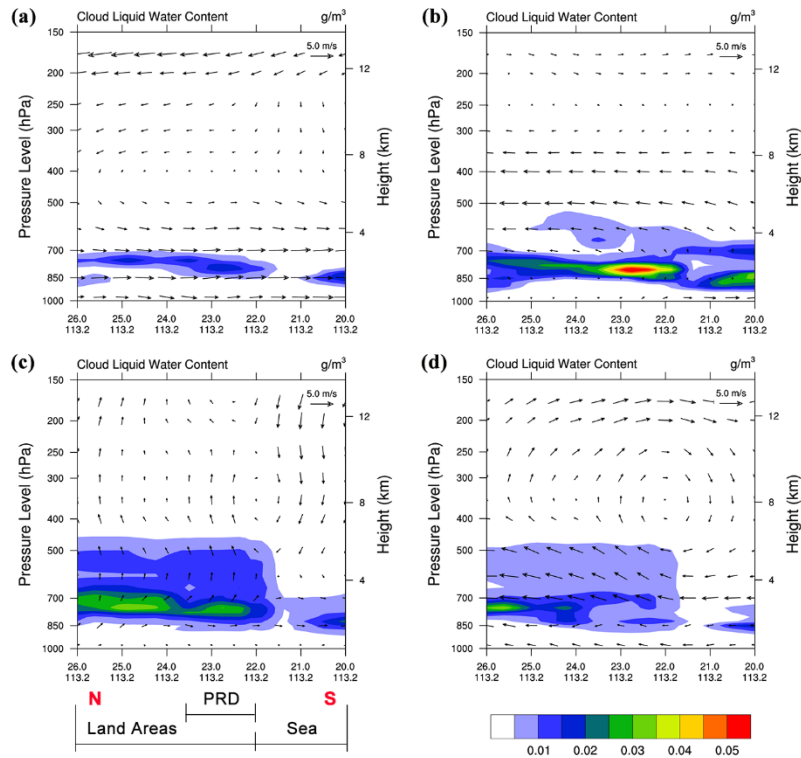
334 3.3 O₃ production conditions: comparison of clouds

335 Clouds efficiently reflect solar radiation (Liou, 1976), and therefore, they have a notable impact on the local formation of O₃.
 336 Figure 6 displays the cross sections of mean ERA-Interim cloud liquid water contents (CLWC) at 14:00 LT of all O₃ pollution
 337 days corresponding to the typhoon-induced and no-typhoon scenarios of two seasons, which were also made along the 113.2 E
 338 longitude line, from 26.0 N to 20.0 N (Fig. S4). The comparison of ~~cloud liquid water content~~ CLWC in the cross sections
 339 ~~(Fig. 6, derived from the ERA-Interim datasets)~~ suggests that typhoons generally resulted in fewer clouds in autumn but more
 340 clouds in summer, which agrees well with the comparison of cloud covers in Table 2. The presence of fewer clouds on the
 341 typhoon-induced days in autumn can be attributed to two reasons: the influence of dry air masses (indicated by lower RH in
 342 Table 2 and Fig. S6) and/or the hindrance of cloud formation by downdrafts. In summer, the strengthened updrafts above the
 343 land caused by typhoons favoured cloud formation, which is demonstrated by higher ~~cloud liquid water content~~ CLWC at the
 344 heights of ~~2–5 km, 500–850 hPa~~ and increases in medium and high cloud covers. In areas above the PRD below 850 hPa,
 345 downdrafts led to slight decrease of clouds in the typhoon-induced scenario in summer, which is also indicated by reduced low
 346 cloud cover. As a consequence of varied cloud covers in each scenario, on average, net surface solar radiation increased by

347 13% and decreased by 7% on the typhoon-induced days in autumn and summer, respectively (Table 2), which promoted and
348 hindered O₃ production in the PRD during these two seasons, respectively.
349



351 **Figure 5.** O_3 mixing ratio (ppbv) and wind fields at the height of (a–d) 850 hPa, (e–h) 700 hPa, and (i–l) 500 hPa at 14:00 LT for ~~the four scenario~~ **scenario**:
 352 (a, e, i) autumn, typhoon-induced; (b, f, j) autumn, no-typhoon; (c, g, k) summer, typhoon-induced; and (d, h, l) summer, no-typhoon. The red triangle in each plot
 353 indicates the PRD. The gridded areas indicate that vertical wind speed is less than 0, or downdrafts occur.



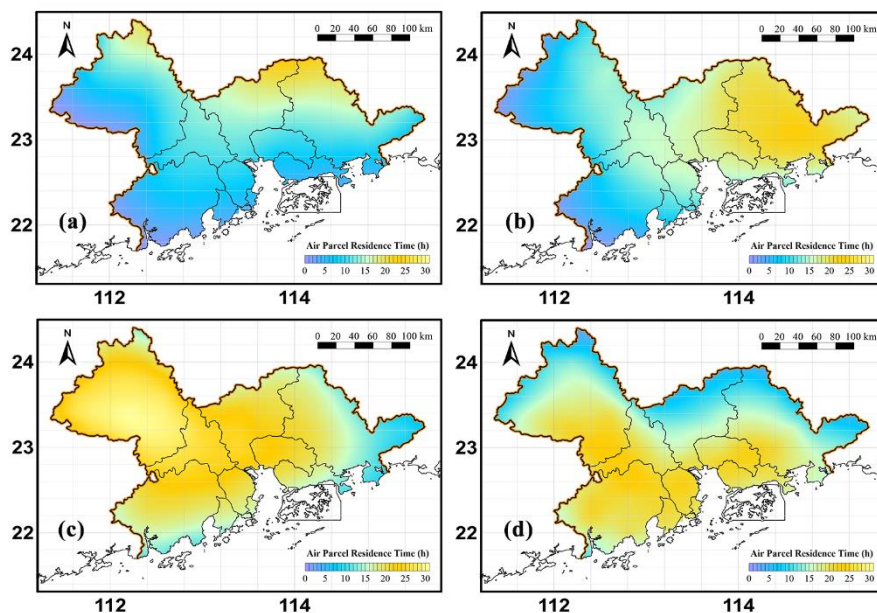
354

355 **Figure 6.** The cross sections of mean cloud liquid water content (g/m^3) and wind vectors at 14:00 LT for ~~the four scenarios~~ each scenario:
 356 (a) autumn, typhoon-induced; (b) autumn, no-typhoon; (c) summer, typhoon-induced; and (d) summer, no-typhoon. Cross sections were
 357 made along the 113.2°E longitude line, ~~are~~ from 26.0°N to 20.0°N along the 113.2°E longitude line (Fig. S4). The vectors indicate meridional
 358 wind speed (m/s) and vertical wind speed (cm/s). PRD, the Pearl River Delta.

359 3.4 O₃ accumulation conditions: comparison of APRTs

360 The longer APRTs are, the more likely that O₃ produced by local emissions accumulates within the targeted region and notably
 361 contributes to near-ground O₃ pollution. In order to study the effect of typhoons on O₃ accumulation, we calculated APRTs in
 362 the PRD in the ~~four~~ typhoon-induced and no-typhoon-scenarios (Fig. 7) for the further comparisons. On the typhoon-induced
 363 days in autumn, APRTs were typically 5–10 hours (mean = 9.5 hours) — shorter than those on the no-typhoon days in autumn
 364 (mean = 13.1 hours). In addition, lower APRT values occurred in the central part of the PRD, where high anthropogenic
 365 emissions of pollutants are distributed (Zheng et al., 2009). Despite more active O₃ chemistry discussed in the last section,
 366 locally sourced O₃ was less likely to accumulate within the PRD in this scenario, potentially limiting the contribution of local
 367 emissions for O₃. The comparison suggests opposite results in the summer scenarios, that is, APRTs on the typhoon-induced
 368 days (20–30 hours, mean = 21.0 hours) were overall higher than those on the no-typhoon days (15–25 hours, mean = 16.5
 369 hours). This favoured the accumulation of locally sourced O₃, and, to some extent, offset the influence of weakened O₃
 370 formation ~~to some extent to maintain high contributions of local emissions to O₃ pollution.~~ In both seasons, Based on the
 371 comparison of O₃ production conditions in the previous section and the comparison of O₃ accumulation conditions in this

372 section, typhoons did not ~~cause~~ provide more favourable conditions for O₃ production and accumulation simultaneously in the
373 PRD in both autumn and summer, thus potentially resulting in a less important role of local contributions in O₃ pollution here.
374 More quantitative evaluations of the contributions from multiple O₃ sources are discussed in Sect. 4.
375



376

377 **Figure 7.** The spatial distributions of APRTs in the PRD for the four scenarios each scenario: (a) autumn, typhoon-induced; (b) autumn, no-
378 typhoon; (c) summer, typhoon-induced; and (d) summer, no-typhoon.

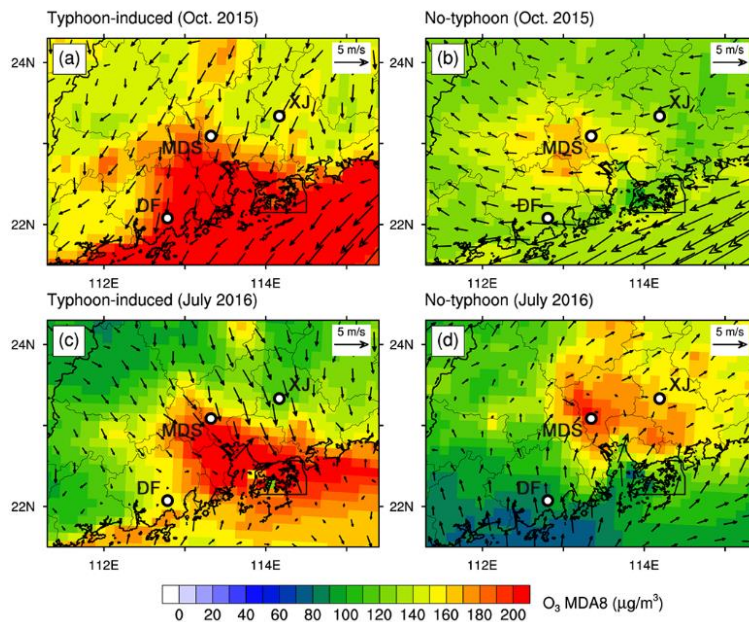
379 3.5 Meteorological conditions on the close typhoon-induced days

380 On the close typhoon-induced days in the two seasons, stronger north winds prevailed and total cloud cover was higher than
381 that on the no-typhoon days (Table 2), suggesting better conditions for the transport of O₃ but less favourable conditions for
382 O₃ production. As displayed in Fig. S7, the APRT values were significantly lower on the close typhoon-induced days (mean
383 = 6.6 hours, 12.9 hours in autumn and summer, respectively) than on the no-typhoon days, making it even harder for locally
384 sourced O₃ to accumulate within the PRD. Therefore, close typhoons are concluded to promote the transport of O₃ from the
385 outside and to reduce the contributions of O₃ produced from local emissions in a more notable way. In addition, close typhoons
386 led to stronger downdrafts in autumn and updrafts in summer than other scenarios in the same season (Fig. S8). It should be
387 noted that the structure of the summer monsoon cell near the PRD was destroyed in the close typhoon-induced scenario in
388 summer, indicating the stronger influence of typhoons on regional wind fields. The dominant role of O₃ transport during O₃
389 pollution days in this special scenario agrees well with the reported episode-based analyses (Lam et al., 2005; Li, 2013).

390 4 Comparisons of O₃ processes and sources

391 The comparisons of meteorological conditions served as qualitative evidence to determine the general influence of typhoons
392 on O₃ transport, production and accumulation in autumn and summer. Based on the comparison between the CMAQ modelling
393 results on typical O₃ pollution days in October 2015 and July 2016, more quantitative evidence can be presented. Figure 8
394 displays modelled mean O₃ MDA8 concentrations and wind fields (at 14:00 LT) in the four scenarios on the typhoon-induced
395 and no-typhoon O₃ pollution days of two seasons. Large standard-exceedance (> 160 µg/m³) areas were distributed in the PRD
396 on most days, and the typhoon-induced days of both seasons generally featured higher O₃ levels. The distinct wind fields for
397 these scenarios, which were consistent with those in the longer timespan (Fig. S5), indeed led to different spatial distributions
398 of O₃. Generally, the most severe O₃ pollution occurred in the downwind areas, such as the central and southern parts of the
399 PRD on the typhoon-induced days in October 2015, the central PRD on the no-typhoon days in October 2015, and the northern
400 and eastern PRD on the no-typhoon days in July 2016. On the typhoon-induced days in July 2016, high levels of O₃
401 accumulated around the PRE. In this section, we discuss the different contributions of various O₃ processes and sources on
402 these days to better understand the effect of typhoons on O₃ pollution in the PRD.

403



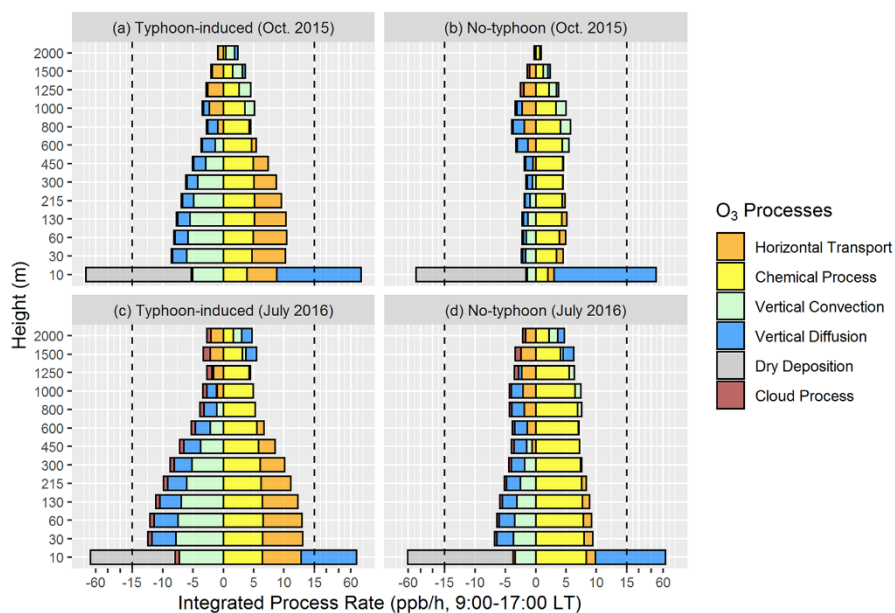
404

405 **Figure 8.** Modelling mean O₃ MDA8 concentrations (µg/m³) and wind vectors (at 14:00 LT) on the representative O₃ pollution days: (a)
406 the typhoon-induced days in October 2015 (14–16 and 21 October 2015); (b) the no-typhoon days in October 2015 (28 October and 3–5
407 November 2015); (c) the typhoon-induced days in July 2016 (7–8 and 30–31 July 2016); and (d) the no-typhoon days in July 2016 (22–26
408 and 29 July 2016). Three representative sites in the PRD are shown as black circles in the plots: XJ, Xijiao; MDS, Modiesha; DF,
409 Duanfen.

410 4.1 O₃ processes: transport vs chemical process

411 The PA tool in CMAQ was used to quantify the contributions of transport and chemical process to the O₃ variations on O₃
412 pollution days under various scenarios in October 2015 and July 2016. As is shown in Fig. 9, the daytime (9:00–17:00 LT) O₃
413 PA results within the PRD in all scenarios share similar characteristics. Dry deposition dominated O₃ removal near the surface,
414 and it also led to high gradients of O₃ concentrations that promote downward O₃ diffusion. Within the PBL (about 0–1 km in
415 height), O₃ was mainly contributed by horizontal transport and chemical process, and vertical convection led to the drop of O₃
416 concentrations. However, differences existed between the O₃ PA results in four the typhoon-induced and no-typhoon scenarios,
417 indicating the impact of typhoons on the transport and production of O₃. In both months, typhoons led to notably higher
418 contribution of horizontal transport to O₃, especially in the lower and middle part of the PBL. Within the PBL, on average, it
419 increased from -0.9 ppb/h, -0.8 ppb/h to 1.2 ppb/h, 2.0 ppb/h under typhoon influence in autumn and summer, respectively.
420 The comparison of the contribution of chemical process (in absolute rates) suggests that they had opposite effects in the two
421 months — under typhoons, the contribution increased in October 2015 (from 4.0 ppb/h to 4.5 ppb/h within the PBL, or by
422 11.4%), but it decreased in July 2016 (from 7.1 ppb/h to 5.7 ppb/h within the PBL, or by -20.8%). In other words, typhoons
423 promoted and hindered O₃ production in autumn and summer, respectively. These results agree well with the comparisons of
424 O₃ transport and production conditions in the previous section.

425



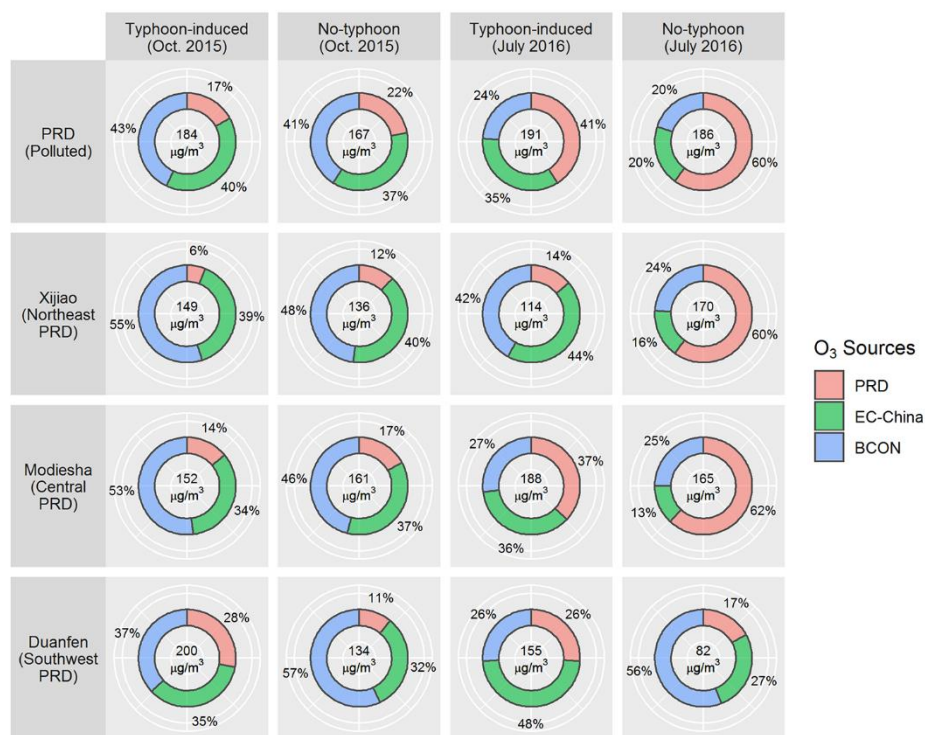
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427 **Figure 9.** The daytime-mean (9:00–17:00 LT) hourly contributions of O₃ processes within the PRD in vertical layers 1–13 on representative
428 O₃ pollution days: (a) the typhoon-induced days in October 2015 (14–16 and 21 October 2015); (b) the no-typhoon days in October 2015
429 (28 October and 3–5 November 2015); (c) the typhoon-induced days in July 2016 (7–8 and 30–31 July 2016); and (d) the no-typhoon days
430 in July 2016 (22–26 and 29 July 2016).

431 4.2 O₃ sources: local sources vs regional sources

432 The contributions of various sources to O₃ within the PRD are determined by the combined impact of O₃ transport, production
 433 and accumulation. The results for ~~the-mean~~ daytime (9:00–17:00 LT) O₃ SA near the ground (about 0–80 m in height) ~~in four~~
 434 ~~scenarios on typhoon-induced and no-typhoon O₃ pollution days~~ are illustrated in Fig. 10. For polluted regions within the PRD,
 435 stronger O₃ production under typhoons did not lead to a higher proportion of local contributions to O₃ pollution in October
 436 2015 — it even decreased from 22% (on the no-typhoon days) to 17% (on the typhoon-induced days). The contributions of
 437 EC-China emissions and BCON, in contrast, increased slightly from 37%, 41% to 40%, 43%, respectively. The distinction of
 438 the O₃ SA results is more apparent for the summer scenarios, that is, typhoons resulted in growing contributions from O₃
 439 transported from other regions (from 40% to 59%) but decreased local contributions (from 60% to 41%) in July 2016. More
 440 favourable O₃ accumulation conditions (indicated by higher APRTs on the representative typhoon-induced O₃ pollution days
 441 in summer (Fig. S9)) were far from sufficient to compensate for the effect of weakened O₃ production on the high contributions
 442 of local sources.

443



444

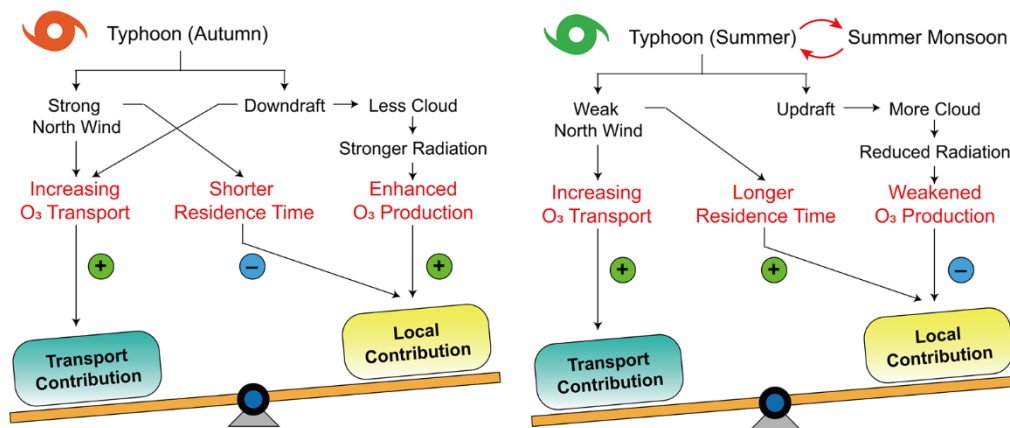
445 **Figure 10.** The mean O₃ SA near the ground (about 0–80 m in height) on ~~representative~~ the represented typhoon-induced and no-typhoon
 446 O₃ pollution days ~~for the four scenarios in October 2015 and July 2016~~ (the average results of 9:00–17:00 LT). The locations of the three
 447 representative sites (Xijiao, Modiesha and Duanfen) are shown in Fig. 8. PRD, the Pearl River Delta; EC-China, East China and Central
 448 China; BCON, the boundary conditions of the d02 modelling.

449 Furthermore, owing to the variations of wind fields, the comparison results of O₃ SA in different parts of the PRD may differ
450 from the regional ones. For instance, while the comparisons of O₃ SA in the Xijiao and Modiesha site (located in the northeast
451 and central part of the PRD, respectively) agree well with those in the polluted regions of the PRD, higher contributions of
452 PRD emissions for O₃ can be found in the Duanfen site (located in the southwest part of the PRD) on the typhoon-induced
453 days of two months in comparison to these on the corresponding no-typhoon days (Fig. 10). Since the site was located in the
454 downwind region in the typhoon-induced scenario in October 2015 (Fig. 8a), enhanced O₃ production led by typhoons from
455 the massive emissions of O₃ precursors in the central PRD (Zheng et al., 2009) contributed to higher local contributions for O₃
456 pollution here ([as the distribution of local contributions in percentage to daytime O₃ shown in Fig. S10](#), the highest local
457 contribution in the PRD occurred in areas near the Duanfen site and almost reached 40% in this scenario, which was even
458 higher than that in the corresponding no-typhoon scenario (33%)). In the no-typhoon scenario in July 2016, the site was located
459 in the upwind regions under the prevailing of southwest winds, limiting the contributions of local emissions for O₃ at the site
460 (Fig. 8d). Thus, higher local contributions can also be found in the typhoon-induced scenario in this month.

461 **5 Discussion and conclusions**

462 The significance of typhoons on O₃ pollution in the PRD calls for thorough evaluations of the different causes of O₃ pollution
463 with the appearance of typhoons in the Northwest Pacific. In this study, we revealed the different impacts of typhoons on O₃
464 transport, production and accumulation in the PRD (as summarised in Fig. 11) through systematic comparisons of
465 meteorological conditions, the contributions of various O₃ processes and sources in the typhoon-induced and no-typhoon
466 scenarios. We found that typhoons tended to promote O₃ transport towards the PRD, but failed to provide more favourable O₃
467 production and accumulation conditions simultaneously, which limited the contribution of local emissions to O₃ pollution.
468 Furthermore, there were also differences between the influence of typhoons on O₃ pollution in autumn and summer. More
469 favourable transport conditions occurred in the typhoon-induced scenario in autumn, which was characterised by higher wind
470 speeds and the increased influence of downdrafts. In summer, the mixed types of air masses in the typhoon-induced scenario
471 were likely to bring more O₃ into the PRD than the clean marine air masses in the no-typhoon scenario, also suggesting
472 enhanced O₃ transport under the influence of typhoons. Generally, typhoons led to cloudless conditions, stronger solar radiation,
473 and thus more rapid O₃ production in autumn, but shorter APRTs (5–10 hours) suggest that locally sourced O₃ was hard to
474 accumulate within the PRD. As a result, the contributions in percentage of local emissions to O₃ pollution decreased (slightly
475 by ~5% for the polluted regions of the PRD in October 2015). In contrast, in summer, intensified updrafts associated with
476 typhoons strengthened cloud formation, weakened solar radiation, and thus restrained local O₃ production. Longer APRTs (>
477 20 hour) under typhoon influence were far from sufficient to maintain high contributions of local emissions for O₃ pollution
478 (which decreased by ~20% for the polluted regions of the PRD in July 2016). However, due to the variations of wind fields
479 under different scenarios, the changes of local and transport contributions for O₃ led by typhoons were different in the
480 southwest part of the PRD, that is, higher contribution from emissions within the PRD and reduced transport contribution

481 occurred in the typhoon-induced scenarios in both seasons. As for the close typhoon-induced scenario, O₃ transport was further
 482 strengthened, but meteorological conditions in the PRD became less favourable for both the production and accumulation of
 483 O₃.
 484



485
 486 **Figure 11.** The summary of the causes of O₃ pollution in the PRD under typhoon influence in autumn and summer.

487 The East Asian monsoon, changing with seasons, has a pronounced impact on local meteorological conditions as well as the
 488 characteristics of O₃ pollution in East China (He et al., 2008). The seasonal behaviour of the East Asian monsoon is likely to
 489 result in the seasonally varied effect of typhoons on O₃ pollution in the PRD. In October, the summer monsoon has almost
 490 finished its retraction and the winter monsoon is beginning (Ding, 1994). Thus, there are not many obstacles to the southward
 491 movement of typhoon periphery and the transport of O₃ towards the PRD by the continental and coastal air masses. Large
 492 downdraft-influenced areas in Central and South China occur in this scenario, and high O₃ levels and low RH in these areas
 493 indicate the potentially important role of directly downward O₃ transport. In July, the summer monsoon reaches its strongest
 494 (Ding, 1994). The interaction between typhoon periphery and the summer monsoon results in stagnation and enhanced updrafts
 495 above the land areas of the PRD and its surroundings. Only when typhoon is close enough to the PRD is the stagnation
 496 terminated and the structure of the summer monsoon cell broken. This also explains why some summertime typhoon-induced
 497 O₃ episodes in the PRD can be typically divided into two periods, as stagnation leads to the accumulation of locally produced
 498 O₃ in the first phase and strong northerly winds strengthen O₃ transport before the landing of typhoons (Lam et al., 2005; Li,
 499 2013). It should be noted that updrafts, rather than downdrafts, prevailed on the typhoon-induced O₃ pollution days in summer.
 500 High levels of O₃ did not necessarily occur in the regions dominated by downdrafts in this scenario, suggesting a less notable
 501 connection between downdrafts and summertime O₃ pollution in the PRD. Further investigations are required to trace the
 502 detailed process of downward O₃ transport, including the stratosphere-troposphere exchange (Stohl et al., 2003), in each
 503 scenario, and quantify their contributions to near-ground O₃ pollution.

504

505 Some limitations remain in this study. We chose O₃ pollution days as individual samples, ignoring the influence of O₃ pollution
506 on the previous days. Thus, more detailed full-episode analyses are required. Moreover, owing to the small sampling size, the
507 influence of typhoons on O₃ pollution in the PRD is still not fully understood, including, for instance, the detailed connections
508 between the features of typhoons (intensity, position) and O₃ pollution. However, the comparisons of meteorological conditions,
509 O₃ processes and sources in different scenarios and seasons demonstrate the complex causes of typhoon-induced O₃ pollution
510 in the PRD — typhoons tend to enhance O₃ transport into the PRD in both seasons, but their impacts on the production and
511 accumulation of O₃ are completely different. As a result, emissions within (outside of) the PRD are likely to contribute less
512 (more) on the typhoon-induced O₃ pollution days than on the no-typhoon days, ~~and more attention should be paid to controlling~~
513 ~~anthropogenic emissions of O₃ precursors on a larger scale under typhoon influence. In order to effectively alleviate O₃~~
514 ~~pollution and to reduce the population exposure in the PRD, more attention should be paid to controlling anthropogenic~~
515 ~~emissions of O₃ precursors on a larger scale, rather than focusing on local emission, under typhoon influence. For air quality~~
516 ~~management, it is suggested to comprehensively evaluate the efficiency of fractional local and non-local emission reductions~~
517 ~~to reduce O₃ levels in the PRD in different scenarios (Thunis et al., 2019; Thunis et al., 2020).~~ This study also suggests that a
518 thorough evaluation of O₃ transport, production and accumulation conditions can be applied to understand the causes of
519 regional O₃ pollution not only in the PRD, but also in other regions. The results will help find efficient strategies to alleviate
520 regional O₃ pollution as well as to reduce its adverse effects.

521
522 *Data availability.* Data are available from the corresponding author upon request.

523
524 *Author contributions.* KQ, XW and YZ designed the study. KQ, XW, and TX did the simulation work, including the operation
525 of the WRF, SMOKE and CMAQ models. JS, HD, LZ and YZ provided observational results of field campaigns and the
526 routine monitoring datasets for the evaluation of model performance. KQ, XW, YY and YZ analysed the modelling results.
527 KQ, XW, YY and YZ wrote and revised this paper, with critical feedbacks from all other authors.

528
529 *Competing interests.* The authors declare no conflict of interest.

530
531 *Acknowledgements.* This work was supported by the National Key Research and Development Program of China (Grant No.
532 2018YFC0213204, 2018YFC0213506, [2018YFC0213501](#)) and the National Science and Technology Pillar Program of China
533 (Grant No. 2014BAC21B01).

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