



- ¹ Time dependent source apportionment of submicron organic
- ² aerosol for a rural site in an alpine valley using a rolling PMF
- 3 window
- 6 Urs Baltensperger¹, Jay G. Slowik¹, Imad El Haddad¹, and André S.H. Prévôt^{1**}
- ¹Laboratory of Atmospheric Chemistry, Paul Scherrer Institute, CH-5232 Villigen PSI,
 8 Switzerland
- 9 ²Datalystica Ltd., Park innovAARE, CH-5234 Villigen, Switzerland
- ³Empa, Swiss Federal Laboratories for Materials Science and Technology, Laboratory for Air
- 11 Pollution and Environmental Technology, CH-8600 Dübendorf, Switzerland
- 12 * G.C. and Y.S. contributed equally to this manuscript
- 13 ** Correspondence to: André S. H. Prévôt (andre.prevot@psi.ch)





15 Abstract

16 We have collected one year of aerosol chemical speciation monitor (ACSM) data in Magadino, 17 a village located in the south of the Swiss Alpine region, which is one of the most polluted 18 areas in Switzerland. We analysed the mass spectra of organic aerosol (OA) by positive matrix 19 factorization (PMF) using Source Finder Professional (SoFi Pro) to retrieve the origins of OA. 20 Therein, we deployed the rolling algorithm to account for the temporal changes of the source 21 profiles, which is closer to the real world. As the first ever application of rolling PMF analysis 22 for a rural cite, we resolved two primary OA factors (traffic-related hydrocarbon-like OA 23 (HOA) and biomass burning OA (BBOA)), one local OA (LOA) factor, a less oxidized 24 oxygenated OA (LO-OOA) factor, and a more oxidized oxygenated OA (MO-OOA) factor. 25 HOA showed stable contributions to the total OA through the whole year ranging from 8.1– 10.1%, while the contribution of BBOA showed a clear seasonal variation with a range of 8.3-26 27 27.4% (highest during winter, lowest during summer) and a yearly average of 17.1%. The OOA 28 was represented by two factors (LO-OOA and MO-OOA) throughout the year. OOA 29 contributed 71.6% of the OA mass, varying from 62.5% (in winter) to 78% (in spring and 30 summer). The uncertainties (σ) for the modelled OA factors (i.e., rotational uncertainty and 31 statistical variability of the sources) varied from $\pm 4\%$ (LOA) to a maximum of $\pm 40\%$ (LO-32 OOA). Considering the fact that BBOA and LO-OOA (showing influences of biomass burning in winter) had significant contributions to the total OA mass, we suggest a reduction and control 33 34 of the residential heating as a mitigation strategy for better air quality and lower PM levels in 35 this region. In Appendix A, we conducted a head-to-head comparison between the conventional 36 seasonal PMF analysis and the rolling mechanism. It showed similar or slightly improved 37 results in terms of mass concentrations, correlations with external tracers and factor profiles of 38 the constrained POA factors. The rolling results show smaller scaled residuals and enhanced 39 correlations between OOA factors and corresponding inorganic salts than those of the seasonal





40 solutions, was most likely because the rolling PMF analysis can capture the temporal variations 41 of the oxidation processes for OOA sources. Specifically, the time dependent factor profiles of 42 MO-OOA and LO-OOA can well explain the temporal viabilities of two main ions for OOA 43 factors, m/z 44 (CO₂⁺) and m/z 43 (mostly C₂H₃O⁺). This rolling PMF analysis therefore 44 provides a more realistic source apportionment (SA) solution, with time-dependent OA sources. 45 The rolling results show also good agreement with offline Aerodyne aerosol mass spectrometer 46 (AMS) SA results from filter samples, except for winter. This is likely because the online 47 measurement is capable of capturing the fast oxidation processes of biomass burning sources. 48 This study demonstrates the strengths of the rolling mechanism and provides a comprehensive 49 criterion list for ACSM users to obtain reproducible SA results and is a role model for similar 50 analyses of such world-wide available data.

51 **1 Introduction**

52 Atmospheric particulate matter (PM) affects human health and climate. In particular, it influences the radiative balance (IPCC, 2014; von Schneidemesser et al., 2015), reduces 53 54 visibility (Chow et al., 2002; Horvath, 1993), and negatively affects human health by triggering 55 respiratory and cardiovascular diseases and allergies (Daellenbach et al., 2020; Dockery and Pope, 1994; Mauderly and Chow, 2008; Monn, 2001; Pope and Dockery, 2006; von 56 57 Schneidemesser et al., 2015). Fine PM exposure strongly correlates with the global mortality 58 rate. Lelieveld et al. (2015) estimated that outdoor air pollution, mostly PM_{2.5} (PM with an 59 aerodynamic diameter smaller than 2.5 µm) causes 3.3 million premature deaths per year 60 worldwide. Despite of this correlation, different aerosol sources may have strongly different 61 effects on health (Daellenbach et al., 2020). Thus, both climate and health effects are affected 62 by particle chemical composition, which is related to emission sources of primary particles and





- 63 precursor gases for secondary aerosol (IPCC, 2014; Jacobson et al., 2000; Jacobson, 2001;
- 64 Lelieveld et al., 2015; Ramanathan et al., 2005).

65 Organic aerosol (OA) constitutes 20–90% fine PM (Jimenez et al., 2009; Murphy et al., 2006; 66 Zhang et al., 2007), and contains millions of chemical compounds. Since OA is subject of an 67 extremely complex mixture of chemical constituents, with highly dynamic spatial and temporal 68 (seasonal, diurnal, etc.) variability of directly emitted particles and gas-phase precursors and a 69 complex chemical processing in the atmosphere, elucidation of the chemical composition and 70 physical properties of OA remains challenging. Identification and quantification of OA sources 71 with a sophisticated interpolation of both spatial and temporal variabilities are essential for a 72 development of effective mitigation strategies for air pollution and a better assessment of the 73 aerosol effect on both health and climate.

74 OA source apportionment (SA) and PM composition has been studied extensively using the 75 Aerodyne aerosol mass spectrometer (AMS) (Canagaratna et al., 2007). However, due to the 76 complexity of the AMS measurements and their high operational expenses, AMS campaigns 77 are often limited to short time periods of a few weeks to months. The aerosol chemical 78 speciation monitor (ACSM) allows for unattended long-term observation (>1 year) of non-79 refractory aerosol particles (Ng et al., 2011a; Fröhlich et al., 2013). It makes it possible to 80 investigate also the long-term temporal variations of OA sources, which is crucial for 81 policymakers to introduce or validate aerosol-related environmental policies.

Positive matrix factorization (PMF) has been used in various studies for SA of OA (Aiken et
al., 2009; Hildebrandt et al., 2011; Lanz et al., 2007; Mohr et al., 2012; Schurman et al., 2015;
Zhang et al., 2011). The multilinear engine (ME-2) implementation of PMF (Paatero, 1999)
improves model performance by allowing the use of *a priori* information (constraints on source
profiles and/or time series) to direct the model towards environmentally meaningful solutions





87 (Canonaco et al., 2013; Crippa et al., 2014; Fröhlich et al., 2015; Lanz et al., 2008; Ripoll et 88 al., 2015). For long-term data (one year or more) with high time resolution, the composition of 89 a given source could change considerably due to the meteorological and seasonal variabilities. 90 However, a major limitation of PMF is the assumption of static factor profiles, such that it fails 91 to respond to these temporal changes. Therefore, long-term chemically speciated data have 92 been evaluated monthly or seasonally (Bressi et al., 2016; Canonaco et al., 2015; Minguillón 93 et al., 2015; Petit et al., 2014; Reyes-Villegas et al., 2016; Ripoll et al., 2015) to at least take 94 the seasonal variations into account. To improve analysis of long-term ACSM datasets, a novel 95 approach that utilizes PMF analysis on a smaller time rolling window was first proposed by Parworth et al. (2015) and further refined using ME-2 by Canonaco et al. (2020). The short 96 97 length of the rolling PMF window allows the PMF model to take the temporal variations of the 98 source profiles into account (e.g., biogenic versus domestic burning influences on oxygenated 99 organic aerosol (OOA)), which normally provides a better separation between OA factors. In 100 addition, using this technique together with bootstrap resampling and a random a-value 101 approach allows users to assess the statistical and rotational uncertainties of the PMF results 102 (Canonaco et al., 2020; Tobler et al., 2020).

In this work, we conducted a one year ACSM measurement from September 2013 to October 2014 in Magadino, located in an alpine valley in southern Switzerland. We present a comprehensive analysis of the ACSM dataset measured in Magadino using a novel PMF technique, the "rolling PMF". In addition, we also compare the results of the rolling PMF with the source apportionment of offline AMS filter samples (Vlachou et al., 2018) and conventional seasonal PMF analysis.





109 2 Methodology

110 2.1 Sampling site

Magadino is in a Swiss alpine valley (46°90'37" N, 85°60'2" E, 204 m.a.s.l.), where the 111 112 sampling site located. This site belongs to the Swiss National Air Pollution Monitoring 113 Network (NABEL, https://www.empa.ch/web/s503/nabel). It is around 1.4 km away from the 114 local train station, Cadenazzo, around 7 km away from the Locarno Airport, and nearly 8 km 115 away from the Lake Maggiore. This station is surrounded by agricultural fields within a rural area, which is considered as a rural background site. It can be potentially affected by domestic 116 117 wood burning, adjacent agricultural activity and transit traffic through the valley. The site 118 topography favours quite high PM levels due to stagnant meteorological conditions or 119 boundary layer inversions, especially in winter. The annual average PM_{10} concentration in 120 Magadino exceeded the annual average PM₁₀ limit value for Switzerland (20 µg·m⁻³) for five 121 years out of the period 2007–2016 (Meteotest, 2017; The Swiss Federal Council, 2018).

122 2.2 ACSM measurements

123 In this study, chemical composition and mass loadings of non-refractory constituents of 124 ambient submicron aerosol particles (NR-PM₁) were measured by an Aerodyne quadrupole 125 ACSM (Ng et al., 2011a). The ACSM uses the same sampling and detection technology as the 126 AMS but is simplified and designated for long-term monitoring applications by reducing 127 maintenance frequency, at the cost of lower sensitivity, restriction to integer mass resolution, 128 and no size measurement. Same as for the AMS, sampled submicron particles enter the 129 instrument through a critical orifice (100 μ m I.D.) at a flow rate of 1.4 cm³ s⁻¹ (at 20 °C and 1 130 atm). The sampling flow will pass either through a particle filter or directly into the system 131 using an automated 3-way switching valve, that is switched every ~30 s. The sampled particles 132 are focused by an aerodynamic lens into a narrow beam and impact on a tungsten surface of 133 around 600 °C, where the non-refractory particles vaporize and are subsequently ionized by an





electron impact source (70 eV). The resulting ions are detected by a quadrupole massspectrometer up to a mass to charge ratio m/z = 148 Th. The particle mass spectrum is represented by the difference of the total ambient air signal and the particle-free signal.

137 The quantification of ACSM data requires an estimation of the fraction of NR-PM₁ that 138 bounces off the oven without being vaporized and therefore is not detected (Canagaratna et al., 139 2007; Matthew et al., 2008). A collection efficiency (CE) factor is typically introduced to 140 correct for particle bounce, which depends on the particulate water content (Matthew et al., 141 2008), ammonium nitrate mass fraction (ANMF) and acidity (Middlebrook et al., 2012). To 142 eliminate humidity effects on CE, a Nafion membrane dryer (Perma Pure MD) was installed 143 on the sampling inlet. In this study, we compared both, a constant CE of 0.45 and a timedependent CE correction suggested by Middlebrook et al., (2012). It showed that data corrected 144 145 with a constant CE had a better correlation and slope closer to 1 when comparing with the chromatographic SO₄²⁻, NO₃⁻, and Cl⁻ anions (Fig. S1a). In addition, as more than 93.5% data 146 147 have an ANMF smaller than 0.4, only 6.5% of data would be impacted by a time-dependent 148 CE correction, therefore, the ammonium nitrate particles doesn't have significant effects on 149 CE for this dataset. Overall, this dataset agrees with external TEOM measurement of both 150 PM_{2.5} and PM₁₀ daily mass concentrations as shown in Fig S1c with a constant CE value.

The ACSM filament burnt out on 14 April, 2014. This was addressed by switching to the backup filament already installed within the instrument (no venting required). Calibration of the relative ionization efficiencies (RIE) of particulate nitrate, sulphate, and ammonium was conducted using size-selected (300 nm) pure NH₄NO₃ and pure (NH₄)₂SO₄ particles. Calibrations of the relative ionisation efficiency (RIE), m/z scale, and the sampling flow was performed every 2 months. In this study, we used the averaged RIEs for nitrate, sulphate, and ammonium, the exact values are shown in Fig S1.





158 2.3 Complementary measurements

Meteorological data, including temperature, precipitation, wind speed, wind direction, and solar radiation are monitored at the NABEL station. In addition, concentrations of trace gases (SO₂, O₃, NO_x), equivalent black carbon (eBC), and PM₁₀ were measured with a time resolution of 10 minutes. We used an aethalometer (AE 31 model by Magee Scientific Inc.) to measure eBC concentrations. Therefore, we conducted SA of eBC by following Zotter et al. (2017) using Ångstrom exponents for eBC from traffic $\alpha_{tr} = 0.9$ and wood burning $\alpha_{wb} = 1.68$. More details about eBC source apportionment are provided in Section 1 of the SI.

166 2.4 Preparation of the data and error matrices for PMF

167 In this study, we used acsm_local_1610 software (Aerodyne Research Inc.) to prepare the PMF 168 input matrix. In total, this dataset includes 19'708 time points and 67 ions. Of these, CO₂⁺-169 related variables (I_{O+} (m/z = 16), I_{HO+} (m/z = 17), and I_{H2O+} (m/z = 18)) were excluded from the 170 spectral matrix prior to a PMF analysis. They are reinserted into the OA factor mass spectra 171 after the PMF analysis using the ratio from the fragmentation table (Allan et al., 2004); the 172 factor concentrations are likewise adjusted. The measurement error matrix was calculated 173 according to Allan et al. (2003, 2004), with a minimum error considered for the uncertainty of 174 all variables in the data matrix as in Ulbrich et al. (2009). Following the recommendations in 175 Paatero and Hopke (2003) and Ulbrich et al. (2009), the measurement uncertainty for variables 176 (m/z) with a signal-to-noise ratio (S/N) < 2 (weak variables) and S/N < 0.2 (bad variables) were 177 increased by a factor of 2 and 10, respectively. In total, 27 weak ACSM variables were down-178 weighted. Additionally, m/z 12 and 13 were not considered during the PMF analyses, due to 179 being noisy and their overall negative signal. Moreover, m/z 15 is not only very noisy (S/N = 180 0.09), but may be also affected by high biases due to potential interference with air signals.





- 181 2.5 Factor analysis of the organic mass spectra
- 182 PMF has been demonstrated to be a useful tool to retrieve the sources of measured organic
- aerosol mass spectra with a bilinear factor model (Paatero and Tapper, 1994; Ulbrich et al.,
- 184 2009):

185

$$x_{ij} = \sum_{k=1}^{p} g_{ik} \times f_{kj} + e_{ij}$$
(1)

186

where x_{ij} is the mass concentration of the j^{th} mass spectral variable in the time point i^{th} ; g_{ik} is the contribution of the k^{th} factor in the i^{th} time point; f_{kj} is the concentration of the j^{th} mass spectral variable in the k^{th} factor; and e_{ij} is the residual of j^{th} variable of the mass spectra in i^{th} time point. The superscript, p represents the number of factors, which is determined by the user. The cost function of PMF uses least squares algorithm by iteratively minimizing the following quantity Q:

193

$$Q = \sum_{i=1}^{n} \sum_{j=1}^{m} (\frac{e_{ij}}{\sigma_{ij}})^2$$
(2)

194

195 where σ_{ij} is an element in the $n \times m$ matrix of the measurement uncertainties, which 196 corresponds point-by-point to x_{ij} . In addition, we normalized quantity $\frac{Q}{Q_{exp}}$ as a mathematical 197 metric during PMF analysis, where the Q_{exp} is:

198





$$Q_{exp} = (n \times m) - p \times (n + m)$$
⁽³⁾

199

The $\frac{Q}{Q_{exp}}$ supports the user to determine the number of factors required for the model by 200 201 investigating the effects on this quantity of adding/removing a factor. However, PMF itself 202 suffers from rotational ambiguity because of the fact that the object function, Q does not provide unique solutions, that is when $\mathbf{G} \cdot \mathbf{F} = \mathbf{G} \cdot \mathbf{T} \cdot \mathbf{T}^{-1} \cdot \mathbf{F}$, PMF provides a similar value of 203 Q but very different solutions (rotated matrix $\overline{\mathbf{G}} = \mathbf{G} \cdot \mathbf{T}$ (rotated factor time series) and $\overline{\mathbf{F}} =$ 204 $\mathbf{T}^{-1} \cdot \mathbf{F}$ (rotated factor profiles)). Only one of or even none of these rotated solutions may be 205 206 atmospherically relevant. The ME-2 solver (Paatero, 1999) enables theoretically full rotational 207 control over the factor solutions, which is implanted here by imposing constraints via the a-208 value approach on one or more elements of F and/or G (Paatero and Hopke, 2009). The a-value 209 (ranging from 0 to 1) determines how much the resulting factor $(f_{j,solution})$ or time series $(g_{j,solution})$ can vary from the input reference factor $(f_{j,reference})$ or time series $(g_{j,reference})$ 210 211 as shown in Eq. 4a and 4b:

212

$$f_{j,solution} = f_{j,reference} \pm a \cdot f_{j,reference}$$
(4a)

$$g_{j,solution} = g_{j,reference} \pm a \cdot g_{j,reference}$$
 (4b)

213

Previous work using *a*-values has shown to efficiently retrieve environmentally reasonable PMF solutions. This is due to the presence of legitimate *a priori* constraints which decrease the degree of rotational ambiguity (Canonaco et al., 2013, 2020; Crippa et al., 2014; Lanz et al., 2008). Here we configured the ME-2 solver and analysed PMF results using SoFi (Source





- 218 Finder, Datalystica Ltd., Villigen, Switzerland) Pro 6.D interface (Canonaco et al., 2013, 2020),
- 219 developed within the IGOR Pro software (WaveMetrics Inc., Lake Oswego, OR, USA).
- 220 Running PMF over the long-term ACSM datasets assumes that the OA source profiles are static 221 within this time window. This can lead to large errors, since OA chemical fingerprints are 222 expected to vary over time (Paatero et al., 2014). For example, Canonaco et al. (2015) showed 223 that the variability of summer and winter OOA cannot be accurately represented by a single 224 pair of OOA profiles. A common way to reduce the model uncertainty arising from this source 225 is to choose a proper number of OA factors (Sug Park et al., 2000), and then perform a PMF 226 analysis on a subset of measurements to capture temporal features of OA chemical fingerprints. 227 Such characterization of OA sources on a seasonal basis has been demonstrated in a number of studies (Crippa et al., 2014; Lanz et al., 2008; Minguillón et al., 2015; Petit et al., 2014; Ripoll 228 229 et al., 2015; Zhang et al., 2019).
- 230 2.6 Rolling PMF analysis with ME-2

231 In this study, we performed PMF runs with a priori constraints (factor profiles) retrieved from 232 seasonal bootstrap analysis (Section 2.2 in the SI) on a small and rolling window (i.e., 1, 7, 14, 233 and 28 days) that could move across the entire dataset with a step of one day (Canonaco et al., 234 2020; Parworth et al., 2015). In addition, we used the bootstrap re-sampling strategy, which 235 can randomly choose a subset of the original matrix and replicate some of the rows/columns to 236 create a new same-size matrix (Efron, 1979). Here, we combined this rolling PMF analysis 237 with the bootstrap strategy and random *a*-values for constrained factor profiles to estimate the 238 statistical and rotational uncertainties of this PMF analysis. More details of this novel technique 239 is found in Canonaco et al. (2020).





240 2.6.1 Window settings

241 In order to retrieve appropriate constraints, we performed PMF *pre-tests* and bootstrap analysis 242 for different seasons. More details of the steps, settings of these analysis can be found in Section 243 2 of the SI. Here, we constrained primary OA factor profiles (hydrocarbon-like OA factor 244 (HOA) and biomass burning OA (BBOA)) as well as the factor profile of a local factor (LOA) 245 using the *a*-value technique in the rolling PMF analysis. The reference profiles of HOA and 246 BBOA were from the winter bootstrapped PMF solution (Dec, Jan, and Feb) as shown in Fig. 247 S6. With a higher contribution of the biomass burning trace ion m/z 60 in the winter, we expect a more representative and robust BBOA profile from the winter solution than from other 248 249 seasons. The LOA profile was retrieved from the summer bootstrapped PMF solution (Jun, Jul, 250 and Aug) (Fig. S6). To allow the factor profile to adapt itself over time, a random a-value 251 within a range of 0.4 with a step of 0.1 is applied for HOA and BBOA. Canonaco et al. (2020) 252 suggested that an upper a-value of 0.4 is sufficient to cover the temporal variation of OA source 253 profiles. Moreover, due to the uniqueness of the LOA chemical profile, it is tightly constrained 254 with a constant *a*-value of 0.05. The LOA factor appeared only after the filament had been 255 changed (14 April, 2014), and its mass spectrum is dominated by nitrogen-containing 256 fragments (at m/z 58, 84, and 98). The instrument setup thus influenced strongly the sensitivity 257 of these components (likely due to influences of surface ionization). Therefore, this factor was 258 considered in the PMF analysis, but no further interpretation of its potential source will be 259 covered in this manuscript.

In total, we constrained HOA and BBOA factors with random *a*-value (0–0.4, with a step of 0.1), and an exact *a*-value (0.05) for LOA factor in the rolling PMF analysis. There are 25 ($N=5\times5$) possible *a*-value combinations within an individual rolling window. Therefore, 50 PMF iterations for each time window are sufficient to cover all possibilities of the *a*-value combinations. With the rolling window of 50 repeats, each data point (except the data within

12





the first and last time window) will actually have many PMF iterations (i.e., N=length of the 265 266 window×50), where bootstrap resampling and random combinations of constraints is performed. This allows to estimate the statistical and rotational uncertainties of the PMF factors 267 268 (Canonaco et al., 2020). To find the optimum length of the time windows, we tested four 269 different lengths of the time windows (N=1, 7, 14, 28) using the same approaches as in 270 Canonaco et al. (2020). We determined the optimum length of the time window based on the 271 number of missing data points (un-modelled data due to the selection based on the criteria) 272 while applying the same thresholds for the same criteria.

273 2.6.2 Criteria settings

274 Performing a rolling analysis for a one-year data with 50 repeats per window requires several 275 tens of thousands of PMF runs. Manual inspection of all PMF runs is impractical and therefore 276 was replaced by monitoring user-defined criterion scores (Canonaco et al., 2020). In this study, 277 R^2 values of the time series of modelled HOA vs NO_x and eBC_{tr} were used for HOA. The 278 BBOA factor was inspected using the variation of m/z=60 explained by BBOA (Table S1). For 279 these time series based criteria, (criterion 1 to criterion 3 in Table S1), we deployed student t-280 test to minimize subjective judgment while determining the thresholds (more discussions in 281 Section 2.3 of the SI).

Typically, OOA factors are dominated by the signals of f_{43} (C₂H₃O⁺ at m/z = 43) and f_{44} (CO₂⁺ at m/z = 44) that correspond to the less and more oxygenated ion fragments (Canonaco et al., 2015; Ng et al., 2010), where *f* is the fraction of a variable, *i.e.* the intensity $I_{m/z}$ normalized by the sum of the intensities of all organic m/z variables. In this study, we were able to retrieve two OOA factors (i.e., more oxidized OOA (MO-OOA) and less oxidized OOA (LO-OOA)) for the whole year, while MO-OOA can be at either at 4th or 5th position because there are two unconstrained factors. Thus, we used the f_{44} for the 4th factor to sort the unconstrained OOA





289	factors to ensure MO-OOA and LO-OOA sitting on the 4 th and 5 th position, respectively. The
290	details of the sorting scheme can be found in Canonaco et al. (2020). At the same time, we also
291	monitored the f_{43} in LO-OOA and f_{44} in MO-OOA to make sure they are not zero. With this set
292	of criteria, we were able to only select "good" (atmospherically relevant) PMF runs before
293	averaging.

3 Results and discussion

295 3.1 Overview of PM₁ sources in Magadino

296 Considering that the major part of eBC is within PM₁ (Schwarz et al., 2013), we added eBC to 297 the total NR-PM1 from the ACSM to perform a mass closure analysis with PM2.5/PM10 from 298 filters. The gravimetric PM2.5 and PM10 show a high correlation with total estimated PM1 (NR-299 PM₁+eBC) (Fig. S1c). The slopes of the linear fits (± 1 standard deviation) are 1.62 ± 0.05 (R^2 = 0.81, N=79) for PM_{2.5} vs. PM₁ and 1.84 ± 0.03 ($R^2 = 0.67$, N=335) for PM₁₀ vs. PM₁. This 300 301 means that the estimated PM_1 comprised 62% and 54% of the $PM_{2.5}$ and PM_{10} mass, 302 respectively. The daily averages of the inorganic species concentrations measured by the 303 ACSM and those measured on the filters by chromatography show a high correlation, with R^2 = 0.83 for SO₄²⁻, $R^2 = 0.82$ for NO₃⁻ and $R^2 = 0.50$ for Cl⁻, with slopes close to 1 (Fig. S1a). 304 The 2-week average of total ammonium and total nitrate measured by offline AMS technique 305 agree rather well with the ACSM ammonium ($R^2 = 0.47$) and nitrate ($R^2 = 0.79$), as shown in 306 307 the plots in Fig. S1b. The ion balance of particulate ammonium, sulphate and nitrate measured 308 by the ACSM showed that the measured aerosol particles were mostly neutral.

309 The daily average PM₁ components are shown in Fig. 1a, with the annual average PM₁ 310 concentration (including eBC) for the period from September 2013 to October 2014 equal to 311 10.2 μ g m⁻³. In winter, the average PM₁ concentration was highest (13.8 μ g·m⁻³), with OA





- 312 contributing 54% to the total PM1 mass. In summer, the average PM1 mass concentration was
- 313 below $10 \,\mu \text{g} \cdot \text{m}^{-3}$, but the relative contribution of the OA fraction increased to 62%.

314 Seasonally averaged diurnal cycles of NR-PM₁ and of eBC are displayed in Fig. 2. In this study, 315 all the data is based on local time (Central European Time). In fall, spring and summer, the 316 diurnals of these pollutants seem to be mainly affected by the development of the BLH, most 317 of the species show similar diurnal trends for these three seasons. In addition, summer has the 318 highest sulphate concentration, due to the enhanced photochemical production. In winter, air 319 pollutants accumulated during evening and night due to the thermal inversion. In general, eBC 320 and organics have higher levels due to enhanced biomass burning emissions and lower 321 boundary layer height (BLH). We observed distinct midday peaks of organics, sulphate, nitrate, 322 ammonium, chloride, and NOx in the winter. Magadino experienced a series of windless, cold, 323 but sunny periods from December 2013 to January 2014, including such sharp peaks (Fig. S3a). 324 It is interpreted to be due to advection within the shallow boundary layer due to the fact that 325 both primary and secondary pollutants increased simultaneously. Local winds were very low 326 near the ground but likely locally and regionally induced orography influenced winds including 327 vertical diffusion processes were initiated during these times that are difficult track without 328 spatially distributed measurements. Such phenomena were not observed during cloudy, cold, 329 and windless days (Fig. S3b) without thermally induced meteorological processes. Unlike other 330 seasons, the dilution process due to vertical mixing happened only after noon time due to strong 331 inversions during the night and late irradiation of the valley surface in winter.

332 3.2 Seasonal PMF Pre-tests

The automated rolling PMF analysis requires the knowledge of the reference profiles as well as the number of factors. In this section, we present how number of factors were determined based on seasonal PMF *pre-tests*. Initially, unconstrained PMF (3 to 6 factors) was performed





336 separately for the different seasons by following the SA guidelines provided by Crippa et al. 337 (2014). Typically, the HOA profile is characterized by a high contribution of alkyl fragments 338 (e.g. m/z = 43, m/z = 57) and the corresponding alkenyl carbo cations (e.g. m/z = 41, m/z = 55), 339 and the factor profile is relatively consistent over time and different locations. The BBOA 340 profile exhibits significant signals at m/z = 60 and m/z = 73, which are well-known fragments, 341 arising from fragmentation of anhydrous sugars present in biomass-related emissions (Alfarra 342 et al., 2007). For the unconstrained PMF runs, the HOA profile is present throughout the whole 343 year, while the BBOA profile exists for all seasons except in summer. However, as shown in 344 Fig. S4, the measured fraction of m/z = 60 during summer was above the background level of 345 biomass burning-related air masses, 0.3% ±0.06% (Aiken et al., 2009; Cubison et al., 2011; 346 DeCarlo et al., 2008). In addition, the scaled residual at m/z = 60 was decreased when a BBOA 347 factor profile was constrained. Thus, we decided to constrain the BBOA factor for all seasons 348 to potentially capture some local events, such as agricultural and open fires in summer.

349 No evidence for the presence of a cooking-related OA (COA) factor was found based on the 350 seasonal pre-analysis of the key fragments (m/z 55 and m/z 57). It shows no difference in the 351 slope of the absolute mass concentration of m/z 55 vs m/z 57 for different hours of the day (Fig 352 S5a), while different seasons show different slopes (Fig S5b). Therefore, a COA factor was not 353 considered in the PMF model. Moreover, a rapid increase of the measured fraction of m/z = 58, 354 84, and 98 together with m/z 39 (potassium signal) was observed after a filament exchange on 355 14 April, 2014. It is likely that the ACSM's sensitivity towards those ions was changed by the 356 filament exchange. Also, this LOA factor was present for spring, summer, and autumn in 2014 357 in unconstrained PMF runs all the time after the filament change. Therefore, we kept this factor 358 for these three seasons.





359	For the factor(s) with secondary origin, PMF models with different number of factors (3-6)
360	were tested to assess if the oxygenated OA (OOA) factor (with a high contribution of m/z 44
361	that is likely dominated by the CO_2^+ ion, derived from decomposition of carboxylic acids
362	(Duplissy et al., 2011)) is separable without mixing with primary organic aerosol (POA) factors
363	(Fig. S6). We conducted these tests independently for different seasons (autumn 2013, winter,
364	spring, summer, autumn 2014).

365 We analysed winter data first by constraining a HOA factor profile (Crippa et al., 2013) with a tight *a*-value of 0.05. The 3-factor solution (with one OOA factor) showed similarly good 366 367 agreement of HOA and BBOA with the external tracers (NOx, eBCtr, eBCwb) as the 4-factor 368 solution (with two OOA factors). However, the scaled residual of m/z 60 was reduced for the 369 solution with two OOA factors. Moreover, the solution with one OOA factor was not sufficient 370 to explain the variabilities of measured f_{44} vs f_{43} (excluding the primary organic aerosol (POA) 371 factors). For 5- and 6-factor solutions the BBOA and LO-OOA factors started to split. 372 Eventually, we selected the 4-factor solution (HOA, BBOA, MO-OOA, LO-OOA) as the best 373 representation of the winter data.

374 After the bootstrap seasonal PMF runs of winter data (details in Section 2 of the SI), we 375 extracted the HOA and BBOA profiles to use them as the reference factor profiles (Fig. S6) for 376 the pre-tests of other seasons. For the spring, summer, and autumn seasons, 3- to 6-factor PMF 377 solutions were modelled separately for each season by constraining the HOA (a-value=0.1) 378 and BBOA (a-value=0.3) profiles. For the 3-factor solution, we observed an OOA factor with 379 some signals at m/z 58, 84, and 98 which we could not relate to a specific source or process. 380 Also the scaled residuals of variables showed significant levels for these three ions. When we 381 increased the number of OA factors from 3 to 4, a factor dominated by m/z 58, 84, and 98 382 emerged, which we named local organic aerosol (LOA). However, the OOA factor still showed





383 slight signals at m/z 58, 84, and 98. An increase in the number of factors from 4 to 5 did not only result in a decrease in $\frac{Q}{Q_{exp}}$, but also in "clean" OOA factors without mixing with the LOA 384 factor. A further increase in the number of factors did not change $\frac{Q}{Q_{exp}}$ substantially (< 1%), 385 and the sixth factor was a mathematical split of the LOA factor with m/z 58 as the dominating 386 387 variable. Thus, the 5-factor PMF model was chosen as the most appropriate for the spring, 388 summer, and autumn 2014. Note that we did not add the LOA factor for the autumn season in 389 2013 since it appeared only after the filament exchange on 14 April, 2014. This LOA factor was included while running PMF because of the rapid drop of the $\frac{Q}{Q_{\text{ourse}}}$ from 4 to 5 factors in 390 391 the PMF model, but the source of this factor will not be discussed in the manuscript.

392 3.3 Full year rolling PMF analysis

393 Here we present the optimized time window size (14 days) (details of the time window 394 optimization are given in Section 4 of the SI and Fig S10). In total, we considered 53.4% of 395 the PMF runs (11087 out of 20750) with only 11 non-modelled data points. The results of the 396 full-year PMF analysis of the 30-min resolved ACSM data are summarized in Fig. 3. The 397 relative contributions of the OA factors are in addition shown in Fig. 3b. The primary traffic 398 related HOA had very little variation (seasonal averages between 8.1 and 10.1%) throughout 399 the year (Fig. 4). In contrast, BBOA showed a distinct yearly cycle (8.3–27.4%) with a yearly 400 averaged contribution of 17.1%. It increased significantly (to 27.4%) in winter which is typical 401 for Alpine valleys (Szidat et al., 2007). It means that biomass burning was the most important 402 primary OA source during the cold season in Magadino. The eBCwb showed similar trends as 403 the BBOA factor time series during the cold seasons (Fig. 3c). The contribution of LOA 404 remained small before the filament was changed on 14 April, 2014, which is expected because 405 we could not retrieve this factor in seasonal unconstrained PMF runs before April 2014.





406	In this study, we retrieved two OOA factors, LO-OOA and MO-OOA. Total OOA (LO-
407	OOA+MO-OOA) contributed substantially to the total OA mass throughout the whole year
408	with an average contribution of 71.6% (Fig. 3b; Fig. 4). In general, the contribution of OOA to
409	the total OA mass did not vary distinctly over the seasons, but reached a maximum of 90.1%
410	on 12 June, 2014, the day with the highest daily average temperature (30.7 °C).

411 In this work, we did head-to-head comparisons between the bootstrap seasonal solutions and 412 the rolling PMF results (see Fig. A1, Fig. A2, Fig. A3, and Table A1 in the Appendix) in terms 413 of mass concentrations, factor profiles, scaled residuals, and correlations between time series 414 for each factor and corresponding external tracers. We found consistent factor profiles and 415 mass concentrations for the constrained factors (i.e., HOA, BBOA, and LOA), while OOA factors showed quite some differences in both mass concentrations and factor profiles. Rolling 416 417 PMF provided slightly better correlations and smaller scaled residuals, therefore, we consider 418 rolling PMF results to be more environmentally reasonable than those of the seasonal PMF 419 (more details in Appendix A).

420 3.3.1 Optimized OA factors retrieved from a rolling PMF model

421 The primary and secondary OA factors retrieved as an annual mean of all optimized PMF 422 solutions together with their diurnal cycles for all seasons are shown in Fig. 5. Seasonal 423 variations of the OOA factor profiles are demonstrated in Fig. 7 and further discussed in more 424 detail in Section 3.3.2. Note that the primary factors (HOA, BBOA, and LOA) were constrained, 425 where the LOA profile was tightly constrained with an *a*-value of 0.05 due to the uniqueness 426 of its chemical profile. Therefore, only a small variation was allowed for LOA, while the HOA 427 and BBOA model profiles varied more due to looser constraints (Fig. S8). HOA and BBOA 428 have averaged a-values of 0.207, and 0.195, respectively. In addition, they both had good 429 agreement with previous studies (Crippa et al., 2014; Ng et al., 2011b). The probability





distribution function (PDF) of applied *a*-values over time was also investigated (Fig. S8). Most
selected runs chose *a*-values of 0.1–0.3 for HOA and BBOA. The OOA factors show larger
variations in the chemical profiles because these two factors were not constrained due to the
high variability of oxidation processes governing the secondary factors.

434 Due to extensive residential wood combustion combined with winter inversions, the 435 concentrations of BBOA and eBC_{wb} were three times higher at night than at midday. As 436 discussed above, during winter, all of the air pollutants, including all PMF factors peaked 437 concurrently at 10-11 a.m. (local time) due to development of the mixed boundary layer (light 438 blue markers in Fig. 2 for total PM1 and Fig. 5b). In summer, an additional local photochemical 439 production led to an increasing MO-OOA mass during the day (red markers in Fig. 5b), similarly to the diurnal behaviour of sulphate (R^2 =0.63). A night-time increase and a daytime 440 441 decrease of the LO-OOA mass during spring and summer apparently followed condensation 442 and re-evaporation cycles of semi-volatile species, similar to the behaviour of ammonium 443 nitrate. Additionally, nocturnal chemistry of NO3/N2O5 radicals could lead to formation of 444 HNO₃ via N₂O₅ hydrolysis and of organic nitrates via oxidation of VOCs (Brown et al., 2004; 445 Dentener and Crutzen, 1993), thus influencing the diurnal cycles of both particulate nitrate and LO-OOA (with $R^2 = 0.48$ for spring and $R^2 = 0.36$ for summer). 446

In Fig. 6, we also present the diurnal cycles of HOA, eBC_{tr} and NO_x with different patterns for weekdays and weekends. The hourly averages of HOA and eBC_{tr} as well as the NO_x mixing ratio peak during the morning and evening rush hours over the weekdays, while on the weekends there is only an evening pollution increase coinciding with the time when people come back from holidays or night-time leisure activities.





453 3.3.2 f_{44}/f_{43} analysis of secondary OA factors

While m/z 44 is mostly from the fragment of CO₂⁺, a fingerprint of oxygenated species, m/z 43 454 455 can originate from C₂H₃O⁺ (a fingerprint of semi-volatile species) or C₃H₇⁺ (a fingerprint of 456 the primary emissions of hydrocarbon-like species) (Canonaco et al., 2015; Duplissy et al., 457 2011; Ng et al., 2010). Thus, f44 and f43 are often used to identify the oxidation state of the 458 factors, which is important to differentiate the MO-OOA and LO-OOA factors. Under the 459 premise that the POA factors and the LOA factor are all well-resolved, it is important to 460 investigate the relationship between the m/z 44 and m/z 43 signals in the OOA factors to 461 determine whether or not one/two OOA factors are sufficient to explain the dataset. In addition, 462 the shapes of the clouds shown in an f_{44} vs f_{43} plot may also include some source-related 463 information. Figure 7 depicts the relationship between f_{44} and f_{43} of the two modelled OOA 464 factors for different seasons. The yellow cloud of data points represents the measured f44 vs f43 465 after subtracting the m/z 44 and m/z 43 signals contributed by the primary HOA, BBOA and 466 LOA factors. They are colour coded by the total OA mass concentration (data points with OA 467 mass concentration below $2 \mu g \cdot m^{-3}$ are hidden).

468 As shown in Fig. 7a, the data points in Sep-Oct (both in 2013 and 2014) were located on the right side of the triangle presented first by Ng et al. (2010), while the November (2013) data 469 470 points were located within the triangle. In addition, the spring and summer data points (Fig. 7c 471 and Fig. 7d) were all located rather on the right side of the triangle, but the winter points lied 472 within the triangle (Fig. 7b). The data points located within the triangle correspond to the time 473 with lower temperature than those are closer to the right side of the triangle (Fig. S9). This 474 could be explained by the increased biogenic OOA emissions when the temperature was higher, 475 as biogenic OOA tends to be distributed along the right side of the triangle (Canonaco et al., 476 2015; Pfaffenberger et al., 2013). Also, when the temperature decreases, the increased biomass





477 emissions make the OOA points to lie vertically within the triangle (Canonaco et al., 2015;

478 Heringa et al., 2011), which is the case for the winter data (Fig. 7b).

- 479 In July 2014, the rolling PMF LO-OOA moved towards the left side of the plot due to increasing influences from m/z 80, m/z 94 (C₂H₆S₂⁺), m/z 95, and m/z 96 (Fig. S7). Because the 480 481 OA signal of m/z 80 is directly calculated from m/z 94 (Allan et al., 2004), we did not 482 investigate the sources of m/z 80. A potential source of these distinct ions in July is dimethyl 483 disulphide, which shows signals at m/2 94, m/2 95, and m/2 96 (NIST Mass Spectrometry Data 484 Center, 2014). Dimethyl disulphide is widely used in pesticides. Considering that the sampling 485 site is in the middle of farmland, and the diurnal variation of m/z 94 appeared to have peaks 486 during the daytime, we considered the LO-OOA in July to be highly affected by the agricultural activities. However, the static factor profiles of summer LO-OOA from the seasonal summer 487 488 solution had much smaller intensities for m/z 80 and m/z 94 (Fig. S6), which enhanced the 489 scaled residuals for these two variables in the seasonal solutions.
- 490 In winter, LO-OOA (Fig. 9b) was highly affected by biomass burning emissions characterized 491 by the presence of m/z 60, 73 (Alfarra et al., 2007), and the LO-OOA position in the f_{44} vs f_{43} 492 space moved towards the right top direction in the plot due to the increasing biogenic influence 493 as the temperature rose (Fig. 7b, Fig. S9) (Canonaco et al., 2015).
- Figure 7 also highlights the advantages of rolling PMF over seasonal PMF due to its timedependent source profiles. For all the seasons, both seasonal and rolling results show that the linear combinations of OOA factors could properly explain most of the measured OOA points. However, with the static OOA factors for seasonal PMF solutions, it remains difficult to capture the variabilities of some measured data points, while the rolling PMF OOA factors are able to move correspondingly with the temporal changes of the clouds, which moves the factor profiles closer to reality and potentially decreases the scaled residuals significantly (Fig. A3).





- Figure S9 also shows the movements of LO-OOA and MO-OOA factor profiles monthly, where LO-OOA moves towards the right direction as the temperature increases, except for the two light blue squares (June and July) in Fig. S9a. It is clear that temperature plays an important role for the positions of LO-OOA and MO-OOA in the f_{44} vs f_{43} space due to its influences on the OOA sources (biogenic or anthropogenic) as well as the atmospheric processes, which is consistent with previous studies in Zurich (Canonaco et al., 2015).
- 507 3.3.3 Statistical and rotational uncertainties

508 As suggested by Canonaco et al. (2020), combining the bootstrap resampling and the random 509 a-value techniques together with the rolling mechanism, we calculated the standard deviation 510 (σ) and the mean (μ) of the mass concentration for each data point from each OA factor in 511 selected "good" PMF runs. We estimated uncertainty of each OA factor using the slope of the 512 linear fit of σ vs μ . (Fig. 8.). Since the LOA factor was tightly constrained with an *a*-value of 513 0.05, it has the smallest variability (4%). Overall, we found relatively smaller errors of HOA, 514 BBOA, and MO-OOA (i.e., 18%, 14%, and 19%, respectively) and an error of 25% for LO-515 OOA which is comparable with the previous study (Canonaco et al., 2020). The errors for both 516 the MO-OOA and the LO-OOA factor showed some temperature dependence. However, this 517 actually varied with time, and the errors did not significantly change when we separated the 518 dataset into four different temperature groups. Still, data points with higher temperature tend 519 to have larger error for the total OOA than with lower temperature (Fig. 8f). This is because 520 more complex aging processes for OOA factors at high temperature (>20 °C) can cause more 521 variability for the OOA factors.

522 3.3.4 Online vs. offline

523 The mass concentrations for HOA, BBOA, and total OOA were compared with corresponding
524 off-line AMS results (Vlachou et al., 2018) (Fig. S11). Despite some disagreement during





525 winter (BBOA and total OOA), BBOA showed a high correlation -with the offline results for 526 both PM₁₀ and PM_{2.5}, with R^2 of 0.83 and 0.84, respectively. The correlation for total OOA 527 was somehow lower, with R^2 of 0.31 and 0.46 for the offline results of PM₁₀ and PM_{2.5} OOA, 528 respectively. The enhanced OOA concentration for the rolling results during winter season 529 compared to the offline SA results (Fig. 9a), as well as the differences between the rolling 530 results and the offline PM2.5/PM10 results regarding BBOA are most likely due to the fact that 531 the LO-OOA was heavily affected by biomass burning (Fig 9b). The offline results apportioned 532 this biomass burning affected LO-OOA into BBOA, whereas the online ACSM measurements 533 with a higher time resolution were capable to capture the fast oxidation process of biomass 534 burning sources. In addition, the rolling PMF technique enabled the LO-OOA factor profile to 535 adapt to the temporal viabilities of OA sources, so the relatively aged biomass burning related 536 sources was apportioned into LO-OOA during winter time by rolling PMF. The yellow line in 537 Fig. 9a depicts the mass concentration of m/z 60 within LO-OOA, which clearly shows 538 significant enhancements during winter, as well as a good agreement with the LO-OOA time 539 series. HOA did not correlate at all, which may be expected because HOA is not water soluble, 540 and therefore has a very low recovery rate of 0.11 for the offline AMS technique based on the 541 previous study by Daellenbach et al. (2016).

542 **4** Conclusions

In this study, we conducted the first rolling PMF analysis on a 13-month ACSM data collected at a rural site of Switzerland. With the help of the a short rolling PMF time window together with the random *a*-value and bootstrap resampling analysis, we obtained a time dependent SA result with error estimations. Overall, we resolved a comprehensive 5-factor solution with HOA, BBOA, LOA, MO-OOA, and LO-OOA. The contribution of HOA was constant during the year (8.1–10.1%), while BBOA showed a clear seasonal variation (8.3–27.4%), which peaked during winter (due to an increased residential heating source) and contributed least in

24





- summer. OOA was a dominant source throughout the year with a contribution of 71.6% on a yearly average. However, the biomass burning source had a strong influence on LO-OOA formation in winter. Together with BBOA, they make residential heating a considerable source at Magadino during winter. Therefore, a mitigation of residential wood combustion should be considered for a reduction of PM levels in Magadino, especially in winter.
- 555 This manuscript also provided a recommended criterion list (Table S1) as well as a novel way 556 to define thresholds with minimum subjective judgements (student's t-test), which could be a 557 leading example for other SoFi Pro users to conduct rolling PMF. To ensure a good 558 representation of the modelled POA factors and to validate the SA results, we also used the 559 correlations between the PMF factor time series and external data. Both HOA and BBOA agreed well with the corresponding external tracers (NO_x, eBC_{tr}, and eBC_{wb}) for the yearly 560 561 cycles, except summer. This is because the aethalometer model for eBC SA has higher 562 uncertainties with smaller eBC_{wb} mass concentrations. Also, NO_x could originate from multiple sources in this season. Therefore, we used HOA vs. eBC and $EV_{60,BBOA}$ to justify these two 563 factors in summer. The correlation of HOA vs eBC had an R^2 of 0.28, with an $EV_{60,BBOA}$ of 564 565 0.55 in summer. Moreover, the MO-OOA and LO-OOA factors correlated well with inorganic SO₄ and NO₃, respectively. The identified primary and secondary OA factor profiles were 566 567 consistent with the OA factors previously found at a variety of urban, rural, and remote 568 European locations.

This paper assessed the statistical and rotational uncertainties of the PMF solution by combining the bootstrap resampling technique and the random *a*-value approach. It shows relatively small errors for constrained factors compared with a previous study in Zurich (Canonaco et al., 2020), and comparable errors for the OOA factors.





573 We also presented a head-to-head comparison between seasonal PMF solutions and the rolling 574 PMF solution. The POA factors showed good agreement between seasonal and rolling PMF 575 solution, while the OOA factors exhibited greater differences. Overall, the rolling PMF 576 retrieved a somewhat better solution in terms of agreement with external tracers, but much 577 better correlations between the OOA factors and corresponding inorganic salts. In addition, the 578 rolling PMF results provided more realistic results by adapting the temporal variations of OOA 579 factors in the f_{44} vs f_{43} space, which also led to much smaller scaled residuals than for the 580 seasonal PMF. The time series of BBOA and total OOA agreed well with those from offline 581 AMS AS results (Vlachou et al., 2018), except for winter when the fast oxidation processes of 582 biomass burning emissions were not captured by the offline AMS technique.

583 Knowledge of diurnal, seasonal and annual changes in OA sources is essential for interpreting 584 the yearly cycles of OA and defining mitigation strategies for air quality. With the help of more 585 accurate and realistic OA sources together with an estimation of the statistical uncertainty of 586 PMF more constraints can be provided both for climate and air quality models. These improved 587 results are therefore highly valuable for policy makers to solve aerosol-related environmental 588 issues.

589 5 Appendix A: Comparison between seasonal and rolling PMF

590 solutions

The bootstrapped seasonal PMF solutions were compared with the full year rolling PMF results as follows. For each factor, the correlations with external data, the ion intensities in the factor profiles, and the mass concentrations retrieved from the two different source apportionment techniques were compared. The correlations of the factor time series with external data (i.e., NO_x, eBC_{tr}, eBC_{wb}, eBC_{totoal}, SO₄, NO₃, and NH₄) are presented in **Table A1**. The rolling results showed generally slightly better correlations between LO-OOA and NO₃, MO-OOA





- 597 and SO₄, and total OOA with NH₄ than the seasonal PMF results, which is consistent with the
- 598 comparison results from Canonaco et al. (2020). A significant improvement was evident for
- 599 LO-OOA vs NO₃ in spring (with R^2 increasing from 0.02 to 0.48). Concerning the correlations
- 600 of POA factors with external data, rolling results and seasonal showed similar results

Factor	Yearly		SON_2013		DJF		MAM		JJA		SON_2014	
	Seasonal	Rolling										
HOA / NO _x	0.37	0.35	0.52	0.5	0.46	0.47	0.34	0.36	0.15	0.15	0.44	0.42
HOA / eBC _{tr}	0.34	0.33	0.29	0.35	0.41	0.42	0.39	0.31	N/A	N/A	0.38	0.39
HOA / eBC	0.55	0.51	0.79	0.77	0.77	0.73	0.5	0.41	0.29	0.28	0.5	0.47
BBOA / eBC _{wb}	0.82	0.82	0.81	0.79	0.84	0.81	0.67	0.6	N/A	N/A	0.3	0.27
MO-OOA / SO4 ²⁻	0.58	0.49	0.49	0.61	0.52	0.49	0.62	0.66	0.63	0.57	0.43	0.46
LO-OOA / NO3 ⁻	0.11	0.32	0.28	0.42	0.28	0.23	0.02	0.48	0.33	0.36	0.19	0.29
OOA/ NH_4^+	0.46	0.44	0.52	0.55	0.34	0.26	0.73	0.75	0.48	0.47	0.57	0.59

601 **Table A1** Correlation coefficients ($R_{pearson}^2$) between the factor contribution and expected 602 tracers over the year and for individual meteorological seasons (p<0.05).

603

604 As shown in Fig. A1 Comparison of the mass concentrations resulting from rolling PMF 605 and from the seasonal analysis for each factor (colour coded by date and time)., which 606 shows a good agreement for two techniques, except for MO-OOA and LO-OOA. In general, 607 the slope of 1.09 for rolling total OOA vs seasonal OOA suggests a slight underestimation of 608 the OOA contribution by the seasonal PMF solutions, while the slope (<1) for HOA and BBOA 609 suggests that the seasonal PMF solutions overestimate HOA and BBOA. In addition, LOA 610 shows the best agreement between the seasonal and rolling solutions, due to the tight constraint 611 of LOA with an *a*-value of 0.05.

The LO-OOA and MO-OOA factors showed worse agreement than the POA factors for the whole dataset. They had good correlations in each meteorological season, however, with different slopes. For instance, seasonal PMF underestimated LO-OOA in spring and fall 2014, but both seasons showed high correlation with rather narrow scattering. The underestimation of LO-OOA by seasonal PMF was compensated by the overestimation of MO-OOA for these





- 617 two seasons, therefore, the summed OOA still showed a high correlation between rolling and
- 618 seasonal PMF results. This is expected, as the rolling PMF allows the source profiles to adapt
- 619 to temporal variations, while seasonal PMF only has static source profiles.



Fig. A1 Comparison of the mass concentrations resulting from rolling PMF and from the seasonal analysis for each factor (colour coded by date and time).

623

The differences in the major variables of the OOA factors (i.e., m/z 44, 43, and 60) shifted the 624 625 mass concentrations significantly. Therefore, we also compared the factor profiles for both 626 techniques (Fig. A2). For instance, LO-OOA during spring showed higher intensity at m/z 44 627 for the rolling PMF results than for the seasonal PMF results (Fig. A2), which caused the 628 underestimation of LO-OOA for the seasonal PMF in spring. When we averaged the total OOA 629 factor using mass-weighted MO-OOA and LO-OOA factors, rolling PMF yielded higher m/z630 60 for all seasons. As a result, seasonal PMF slightly underestimated the summed OOA factors 631 by around 9%, but slightly overestimated the POA factors by less than <6%.





- The profiles of the constrained factors (HOA, BBOA, LOA) from the rolling results show very
- 633 high correlation with the seasonal results (Fig. A2), which suggests that the primary factors
- and the tightly constrained factor (LOA) were consistent with the static profiles from the
- 635 seasonal PMF analysis.



Fig. A2 Profile comparisons between rolling results and seasonal results for each factor (logscale).

639

We compared the scaled residuals from both source apportionment techniques (**Fig. A3**). The rolling PMF solution had smaller scaled residuals (narrower histogram and the centre was closer to 0) than that of the seasonal PMF solution, which is expected because rolling PMF had more flexibility to adapt to the temporal variabilities of the OA sources.







Fig. A3 Distribution of the scaled residuals over the whole year for the seasonal solution (a)and rolling solution (b).

648

Summarizing, HOA and BBOA were consistent for both rolling and seasonal PMF analysis in terms of the time series, correlations with external tracers, and factor profiles due to the consistency of their chemical factor profiles. In contrast, the MO-OOA and LO-OOA factors were more scattered in terms of averaged factor profiles and mass concentration, which suggests that seasonal PMF analysis was not sufficient to capture these temporal variabilities of their oxidation processes. Also, rolling PMF showed smaller scaled residuals. Therefore, we conclude that the rolling PMF analysis provides more realistic results than the seasonal analysis.

656 **Data Availability**

The data are available upon request to the corresponding author.

658 **Competing interests**

659 The authors declare no competing interests in any form for this work.





660 Author contributions

G. C. and Y. S. contributed equally for this manuscript. G. C. wrote the manuscript, illustrations
as well as data treatments and processing. Y.S. wrote the preliminary manuscript and analysed
preliminary results. R. F. and P. G. helped to run the campaign. P. G., and C. H. provided
external data to validate PMF solution. F.C. provided technique support for SoFi Pro. F.C., A.
T., K. R. D., A. V., J.G.S., I. EI. H., U. B., and A. S. H. P. participated discussions for this
study.

667 Acknowledgements

The ACSM measurements were supported by the Swiss Federal Office for the Environment 668 669 (FOEN). The leading role of the Environmental group of the Swiss Federal Laboratories for Materials and Testing (Empa) in supporting the measurements is very much appreciated. Y. S. 670 acknowledges supports by the "Wiedereinsteigerinnen Program" at the Paul Scherrer Institute. 671 672 This study was also supported by the cost action of Chemical On-Line cOmpoSition and Source Apportionment of fine aerosol (COLOSSAL, CA16109), a COST related project of the Swiss 673 674 National Science Foundation, Source apportionment using long-term Aerosol Mass 675 Spectrometry and Aethalometer Measurements (SAMSAM, IZCOZ0_177063), as well as the 676 EU Horizon 2020 Framework Programme via the ERA-PLANET project SMURBS (grant 677 agreement no. 689443).

678 **References**

Aiken, A. C., Salcedo, D., Cubison, M. J., Huffman, J. A., DeCarlo, P. F., Ulbrich, I. M.,
Docherty, K. S., Sueper, D., Kimmel, J. R., Worsnop, D. R., Trimborn, A., Northway, M.,

- 681 Stone, E. A., Schauer, J. J., Volkamer, R. M., Fortner, E., de Foy, B., Wang, J., Laskin, A.,
- 682 Shutthanandan, V., Zheng, J., Zhang, R., Gaffney, J., Marley, N. A., Paredes-Miranda, G.,





- 683 Arnott, W. P., Molina, L. T., Sosa, G. and Jimenez, J. L.: Mexico City aerosol analysis during
- 684 MILAGRO using high resolution aerosol mass spectrometry at the urban supersite (T0) Part
- 1: Fine particle composition and organic source apportionment, Atmos. Chem. Phys., 9(17),
- 686 6633–6653, doi:10.5194/acp-9-6633-2009, 2009.
- 687 Alfarra, M. R., Prevot, A. S. H., Szidat, S., Sandradewi, J., Weimer, S., Lanz, V. A., Schreiber,
- D., Mohr, M. and Baltensperger, U.: Identification of the Mass Spectral Signature of Organic
 Aerosols from Wood Burning Emissions, Environ. Sci. Technol., 41(16), 5770–5777,
- 690 doi:10.1021/es062289b, 2007.
- 691 Allan, J. D., Alfarra, M. R., Bower, K. N., Williams, P. I., Gallagher, M. W., Jimenez, J. L.,
- 692 McDonald, A. G., Nemitz, E., Canagaratna, M. R., Jayne, J. T., Coe, H. and Worsnop, D. R.:
- 693 Quantitative sampling using an Aerodyne aerosol mass spectrometer 2. Measurements of fine
- 694 particulate chemical composition in two U.K. cities, J. Geophys. Res. Atmos., 108(D3), n/a-
- 695 n/a, doi:10.1029/2002JD002359, 2003.
- 696 Allan, J. D., Delia, A. E., Coe, H., Bower, K. N., Alfarra, M. R. R., Jimenez, J. L., Middlebrook, 697 A. M., Drewnick, F., Onasch, T. B., Canagaratna, M. R., Jayne, J. T. and Worsnop, D. R.: A 698 generalised method for the extraction of chemically resolved mass spectra from Aerodyne aerosol 699 909-922, mass spectrometer data, J. Aerosol Sci., 35(7), 700 doi:10.1016/j.jaerosci.2004.02.007, 2004.
- Bressi, M., Cavalli, F., Belis, C. A., Putaud, J.-P., Fröhlich, R., Martins dos Santos, S., Petralia,
 E., Prévôt, A. S. H., Berico, M., Malaguti, A. and Canonaco, F.: Variations in the chemical
 composition of the submicron aerosol and in the sources of the organic fraction at a regional
 background site of the Po Valley (Italy), Atmos. Chem. Phys., 16(20), 12875–12896,
 doi:10.5194/acp-16-12875-2016, 2016.





- 706 Brown, S. S., Dibb, J. E., Stark, H., Aldener, M., Vozella, M., Whitlow, S., Williams, E. J.,
- 707 Lerner, B. M., Jakoubek, R., Middlebrook, A. M., DeGouw, J. A., Warneke, C., Goldan, P. D.,
- 708 Kuster, W. C., Angevine, W. M., Sueper, D. T., Quinn, P. K., Bates, T. S., Meagher, J. F.,
- 709 Fehsenfeld, F. C. and Ravishankara, A. R.: Nighttime removal of NO x in the summer marine
- 710 boundary layer, Geophys. Res. Lett., 31(7), n/a-n/a, doi:10.1029/2004GL019412, 2004.
- 711 Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q., Onasch,
- 712 T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R., Trimborn, A. M.,
- 713 Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P. and Worsnop, D. R.: Chemical
- 714 and microphysical characterization of ambient aerosols with the aerodyne aerosol mass
- 715 spectrometer, Mass Spectrom. Rev., 26(2), 185–222, doi:10.1002/mas.20115, 2007.
- Canonaco, F., Crippa, M., Slowik, J. G., Baltensperger, U. and Prévôt, A. S. H. H.: SoFi, an
 IGOR-based interface for the efficient use of the generalized multilinear engine (ME-2) for the
 source apportionment: ME-2 application to aerosol mass spectrometer data, Atmos. Meas.
 Tech., 6(12), 3649–3661, doi:10.5194/amt-6-3649-2013, 2013.
- Canonaco, F., Slowik, J. G., Baltensperger, U. and Prévôt, A. S. H.: Seasonal differences in
 oxygenated organic aerosol composition: implications for emissions sources and factor
 analysis, Atmos. Chem. Phys., 15(12), 6993–7002, doi:10.5194/acp-15-6993-2015, 2015.
- 723 Canonaco, F., Tobler, A., Chen, G., Sosedova, Y., Slowik, J. G., Bozzetti, C., Daellenbach,
- 724 Kaspar Rudolf Haddad, I. El, Crippa, M., Huang, R.-J., Furger, M., Baltensperger, U. and
- 725 Prevot, A. S. H.: A new method for long-term source apportionment with time-dependent factor
- 726 profiles and uncertainty assessment using SoFi Pro: application to one year of organic aerosol
- 727 data, Atmos. Meas. Tech. Discuss., doi:10.5194/amt-2020-204, 2020.
- 728 Chow, J. C., Bachmann, J. D., Wierman, S. S. G., Mathai, C. V., Malm, W. C., White, W. H.,





- 729 Mueller, P. K., Kumar, N. and Watson, J. G.: Visibility: Science and Regulation, J. Air Waste
- 730 Manage. Assoc., 52(9), 973–999, doi:10.1080/10473289.2002.10470844, 2002.
- 731 Crippa, M., DeCarlo, P. F., Slowik, J. G., Mohr, C., Heringa, M. F., Chirico, R., Poulain, L.,
- 732 Freutel, F., Sciare, J., Cozic, J., Di Marco, C. F., Elsasser, M., Nicolas, J. B., Marchand, N.,
- 733 Abidi, E., Wiedensohler, A., Drewnick, F., Schneider, J., Borrmann, S., Nemitz, E.,
- 734 Zimmermann, R., Jaffrezo, J.-L., Prévôt, A. S. H. and Baltensperger, U.: Wintertime aerosol
- 735 chemical composition and source apportionment of the organic fraction in the metropolitan
- 736 area of Paris, Atmos. Chem. Phys., 13(2), 961–981, doi:10.5194/acp-13-961-2013, 2013.
- 737 Crippa, M., Canonaco, F., Lanz, V. A., Äijälä, M., Allan, J. D., Carbone, S., Capes, G.,
- 738 Ceburnis, D., Dall'Osto, M., Day, D. A., DeCarlo, P. F., Ehn, M., Eriksson, A.,
- 739 Freney, E., Hildebrandt Ruiz, L., Hillamo, R., Jimenez, J. L., Junninen, H., Kiendler-Scharr,
- 740 A., Kortelainen, A.-M., Kulmala, M., Laaksonen, A., Mensah, A. A., Mohr, C., Nemitz, E.,
- 741 O'Dowd, C., Ovadnevaite, J., Pandis, S. N., Petäjä, T., Poulain, L., Saarikoski, S.,
- 742 Sellegri, K., Swietlicki, E., Tiitta, P., Worsnop, D. R., Baltensperger, U. and Prévôt, A. S. H.
- 743 H.: Organic aerosol components derived from 25 AMS data sets across Europe using a
- consistent ME-2 based source apportionment approach, Atmos. Chem. Phys., 14(12), 6159-
- 745 6176, doi:10.5194/acp-14-6159-2014, 2014.
- 746 Cubison, M. J., Ortega, A. M., Hayes, P. L., Farmer, D. K., Day, D., Lechner, M. J., Brune, W.
- 747 H., Apel, E., Diskin, G. S., Fisher, J. A., Fuelberg, H. E., Hecobian, A., Knapp, D. J., Mikoviny,
- 748 T., Riemer, D., Sachse, G. W., Sessions, W., Weber, R. J., Weinheimer, A. J., Wisthaler, A.
- 749 and Jimenez, J. L.: Effects of aging on organic aerosol from open biomass burning smoke in
- aircraft and laboratory studies, Atmos. Chem. Phys., 11(23), 12049–12064, doi:www.atmos-
- 751 chem-phys.net/11/12049/2011/, 2011.
- 752 Daellenbach, K. R., Bozzetti, C., Křepelová, A., Canonaco, F., Wolf, R., Zotter, P., Fermo, P.,





- 753 Crippa, M., Slowik, J. G., Sosedova, Y., Zhang, Y., Huang, R.-J. J., Poulain, L., Szidat, S.,
- 754 Baltensperger, U., El Haddad, I. and Prévôt, A. S. H. H.: Characterization and source
- 755 apportionment of organic aerosol using offline aerosol mass spectrometry, Atmos. Meas. Tech.,
- 756 9(1), 23–39, doi:10.5194/amt-9-23-2016, 2016.
- 757 Daellenbach, K. R., Uzu, G., Jiang, J., Cassagnes, L.-E., Leni, Z., Vlachou, A., Stefenelli, G.,
- 758 Canonaco, F., Weber, S., Segers, A., Kuenen, J. J. P., Schaap, M., Favez, O., Albinet, A.,
- 759 Aksoyoglu, S., Dommen, J., Baltensperger, U., Geiser, M., El Haddad, I., Jaffrezo, J.-L. and
- 760 Prévôt, A. S. H.: Sources of particulate-matter air pollution and its oxidative potential in Europe,
- 761 Nature, 587(7834), 414–419, doi:10.1038/s41586-020-2902-8, 2020.
- 762 DeCarlo, P. F., Dunlea, E. J., Kimmel, J. R., Aiken, A. C., Sueper, D., Crounse, J., Wennberg,
- 763 P. O., Emmons, L., Shinozuka, Y., Clarke, A., Zhou, J., Tomlinson, J., Collins, D. R., Knapp,
- 764 D., Weinheimer, A. J., Montzka, D. D., Campos, T. and Jimenez, J. L.: Fast airborne aerosol
- 765 size and chemistry measurements above Mexico City and Central Mexico during the
- 766 MILAGRO campaign, Atmos. Chem. Phys., 8(14), 4027–4048, doi:10.5194/acp-8-4027-2008,
- 767 2008.
- 768 Dentener, F. J. and Crutzen, P. J.: Reaction of N 2 O 5 on tropospheric aerosols: Impact on the
- 769 global distributions of NO x , O 3 , and OH, J. Geophys. Res. Atmos., 98(D4), 7149–7163,
- 770 doi:10.1029/92JD02979, 1993.
- 771 Dockery, D. W. and Pope, C. A.: Acute Respiratory Effects of Particulate Air Pollution, Annu.
- 772 Rev. Public Health, 15(1), 107–132, doi:10.1146/annurev.pu.15.050194.000543, 1994.
- 773 Duplissy, J., DeCarlo, P. F., Dommen, J., Alfarra, M. R., Metzger, A., Barmpadimos, I., Prevot,
- A. S. H., Weingartner, E., Tritscher, T., Gysel, M., Aiken, A. C., Jimenez, J. L., Canagaratna,
- 775 M. R., Worsnop, D. R., Collins, D. R., Tomlinson, J. and Baltensperger, U.: Relating





- 776 hygroscopicity and composition of organic aerosol particulate matter, Atmos. Chem. Phys.,
- 777 11(3), 1155–1165, doi:10.5194/acp-11-1155-2011, 2011.
- 778 Fröhlich, R., Cubison, M. J., Slowik, J. G., Bukowiecki, N., Prévôt, A. S. H. H., Baltensperger,
- 779 U., Schneider, J., Kimmel, J. R., Gonin, M., Rohner, U., Worsnop, D. R. and Jayne, J. T.: The
- 780 ToF-ACSM: a portable aerosol chemical speciation monitor with TOFMS detection, Atmos.
- 781 Meas. Tech., 6(11), 3225–3241, doi:10.5194/amt-6-3225-2013, 2013.
- 782 Fröhlich, R., Crenn, V., Setyan, A., Belis, C. A., Canonaco, F., Favez, O., Riffault, V., Slowik,
- 783 J. G., Aas, W., Aijälä, M., Alastuey, A., Artiñano, B., Bonnaire, N., Bozzetti, C., Bressi, M.,
- 784 Carbone, C., Coz, E., Croteau, P. L., Cubison, M. J., Esser-Gietl, J. K., Green, D. C., Gros, V.,
- 785 Heikkinen, L., Herrmann, H., Jayne, J. T., Lunder, C. R., Minguillón, M. C., Močnik, G.,
- 786 O'Dowd, C. D., Ovadnevaite, J., Petralia, E., Poulain, L., Priestman, M., Ripoll, A.,
- 787 Sarda-Estève, R., Wiedensohler, A., Baltensperger, U., Sciare, J. and Prévôt, A. S. H.: ACTRIS
- 788 ACSM intercomparison Part 2: Intercomparison of ME-2 organic source apportionment
- results from 15 individual, co-located aerosol mass spectrometers, Atmos. Meas. Tech., 8(6),
- 790 2555–2576, doi:10.5194/amt-8-2555-2015, 2015.
- Heringa, M. F., DeCarlo, P. F., Chirico, R., Tritscher, T., Dommen, J., Weingartner, E., Richter,
 R., Wehrle, G., Prévôt, A. S. H. and Baltensperger, U.: Investigations of primary and secondary
 particulate matter of different wood combustion appliances with a high-resolution time-offlight aerosol mass spectrometer, Atmos. Chem. Phys., 11(12), 5945–5957, doi:10.5194/acp11-5945-2011, 2011.
- Hildebrandt, L., Kostenidou, E., Lanz, V. A., Prevot, A. S. H. H., Baltensperger, U.,
 Mihalopoulos, N., Laaksonen, A., Donahue, N. M. and Pandis, S. N.: Sources and atmospheric
 processing of organic aerosol in the Mediterranean: insights from aerosol mass spectrometer
 factor analysis, Atmos. Chem. Phys., 11(23), 12499–12515, doi:10.5194/acp-11-12499-2011,





```
800 2011.
```

- 801 Horvath, H.: Atmospheric light absorption-A review, Atmos. Environ. Part A. Gen. Top.,
- 802 27(3), 293–317, doi:10.1016/0960-1686(93)90104-7, 1993.
- 803 IPCC: Clouds and Aerosols, in Climate Change 2013 The Physical Science Basis, edited by
- 804 Intergovernmental Panel on Climate Change, pp. 571–658, Cambridge University Press,
- 805 Cambridge., 2014.
- Jacobson, M. C., Hansson, H.-C., Noone, K. J. and Charlson, R. J.: Organic atmospheric
 aerosols: Review and state of the science, Rev. Geophys., 38(2), 267–294,
 doi:10.1029/1998RG000045, 2000.
- Jacobson, M. Z.: Global direct radiative forcing due to multicomponent anthropogenic and
 natural aerosols, J. Geophys. Res. Atmos., 106(D2), 1551–1568, doi:10.1029/2000JD900514,
 2001.
- 812 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H. H., Zhang, Q., Kroll, J. 813 H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. 814 M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., 815 Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, 816 P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., 817 818 Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., 819 Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., 820 Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., 821 Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., Worsnop, D. R., Dunlea, E. 822 J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y.,





- 823 Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell,
- 824 L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y.
- 825 M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M.,
- 826 Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U. and Worsnop,
- 827 D. R.: Evolution of Organic Aerosols in the Atmosphere, Science (80-.)., 326(5959), 1525-
- 828 1529, doi:10.1126/science.1180353, 2009.
- 829 Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C. and Prévôt, A. S.
- 830 H.: Source apportionment of submicron organic aerosols at an urban site by factor analytical
- modelling of aerosol mass spectra, Atmos. Chem. Phys., 7(6), 1503–1522, doi:10.5194/acp-71503-2007, 2007.
- Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C., Szidat, S., Wehrli,
 M. N., Wacker, L., Weimer, S., Caseiro, A., Puxbaum, H. and Prevot, A. S. H.: Source
 Attribution of Submicron Organic Aerosols during Wintertime Inversions by Advanced Factor
 Analysis of Aerosol Mass Spectra, Environ. Sci. Technol., 42(1), 214–220,
 doi:10.1021/es0707207, 2008.
- Lelieveld, J., Evans, J. S., Fnais, M., Giannadaki, D. and Pozzer, A.: The contribution of
 outdoor air pollution sources to premature mortality on a global scale, Nature, 525(7569), 367–
 371, doi:10.1038/nature15371, 2015.
- Matthew, B. M., Middlebrook, A. M. and Onasch, T. B.: Collection Efficiencies in an
 Aerodyne Aerosol Mass Spectrometer as a Function of Particle Phase for Laboratory Generated
 Aerosols, Aerosol Sci. Technol., 42(11), 884–898, doi:10.1080/02786820802356797, 2008.
- Mauderly, J. L. and Chow, J. C.: Health Effects of Organic Aerosols, Inhal. Toxicol., 20(3),
 257–288, doi:10.1080/08958370701866008, 2008.





- 846 Meteotest: Data Report Switzerland 2007 2016, Bern, Switzerland., 2017.
- 847 Middlebrook, A. M., Bahreini, R., Jimenez, J. L. and Canagaratna, M. R.: Evaluation of
- 848 Composition-Dependent Collection Efficiencies for the Aerodyne Aerosol Mass Spectrometer
- ⁸⁴⁹ using Field Data, Aerosol Sci. Technol., 46(3), 258–271, doi:10.1080/02786826.2011.620041,
- 850 2012.
- Minguillón, M. C., Ripoll, A., Pérez, N., Prévôt, A. S. H., Canonaco, F., Querol, X. and
 Alastuey, A.: Chemical characterization of submicron regional background aerosols in the
 western Mediterranean using an Aerosol Chemical Speciation Monitor, Atmos. Chem. Phys.,
 15(11), 6379–6391, doi:10.5194/acp-15-6379-2015, 2015.
- Mohr, C., Decarlo, P. F., Heringa, M. F., Chirico, R., Slowik, J. G., Richter, R., Reche, C.,
 Alastuey, A., Querol, X., Seco, R., Crippa, M., Zimmermann, R., Baltensperger, U., Barcelona,
 D., Munchen, H. Z. and Mass, J.: Wintertime aerosol chemical composition and source
 apportionment of the organic fraction in the metropolitan area of Paris, Atmos. Chem. Phys.,
 12(4), 1649–1665, doi:10.5194/acp-13-961-2013, 2012.
- Monn, C.: Exposure assessment of air pollutants: a review on spatial heterogeneity and
 indoor/outdoor/personal exposure to suspended particulate matter, nitrogen dioxide and ozone,
 Atmos. Environ., 35(1), 1–32, doi:10.1016/S1352-2310(00)00330-7, 2001.
- 863 Murphy, D. M., Cziczo, D. J., Froyd, K. D., Hudson, P. K., Matthew, B. M., Middlebrook, A.
- M., Peltier, R. E., Sullivan, A., Thomson, D. S. and Weber, R. J.: Single-particle mass
 spectrometry of tropospheric aerosol particles, J. Geophys. Res. Atmos., 111(D23), n/a-n/a,
 doi:10.1029/2006JD007340, 2006.
- 867 Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M., Kroll, J. H.,
- 868 Docherty, K. S., Chhabra, P. S., Bahreini, R., Murphy, S. M., Seinfeld, J. H., Hildebrandt, L.,





- 869 Donahue, N. M., DeCarlo, P. F., Lanz, V. A., Prévôt, A. S. H. H., Dinar, E., Rudich, Y. and
- 870 Worsnop, D. R.: Organic aerosol components observed in Northern Hemispheric datasets from
- 871 Aerosol Mass Spectrometry, Atmos. Chem. Phys., 10(10), 4625–4641, doi:10.5194/acp-10-
- 872 4625-2010, 2010.
- 873 Ng, N. L., Herndon, S. C., Trimborn, A., Canagaratna, M. R., Croteau, P. L., Onasch, T. B.,
- Sueper, D., Worsnop, D. R., Zhang, Q., Sun, Y. L. and Jayne, J. T.: An Aerosol Chemical
 Speciation Monitor (ACSM) for Routine Monitoring of the Composition and Mass
 Concentrations of Ambient Aerosol, Aerosol Sci. Technol., 45(7), 780–794,
 doi:10.1080/02786826.2011.560211, 2011a.
- Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Zhang, Q., Ulbrich, I. M. and Worsnop, D. R.:
 Real-Time Methods for Estimating Organic Component Mass Concentrations from Aerosol
 Mass Spectrometer Data, Environ. Sci. Technol., 45(3), 910–916, doi:10.1021/es102951k,
 2011b.
- NIST Mass Spectrometry Data Center: Disulfide, dimethyl, SRD 69. [online] Available from:
 https://webbook.nist.gov/cgi/cbook.cgi?ID=C624920&Mask=200#Refs (Accessed 6 August
 2020), 2014.
- Paatero, P.: The Multilinear Engine—A Table-Driven, Least Squares Program for Solving
 Multilinear Problems, Including the n -Way Parallel Factor Analysis Model, J. Comput. Graph.
 Stat., 8(4), 854–888, doi:10.1080/10618600.1999.10474853, 1999.
- Paatero, P. and Hopke, P. K.: Discarding or downweighting high-noise variables in factor
 analytic models, Anal. Chim. Acta, 490(1–2), 277–289, doi:10.1016/S0003-2670(02)01643-4,
 2003.
- 891 Paatero, P. and Hopke, P. K.: Rotational tools for factor analytic models, J. Chemom., 23(2),

40





- 892 91–100, doi:10.1002/cem.1197, 2009.
- 893 Paatero, P. and Tapper, U.: Positive matrix factorization: A non-negative factor model with
- 894 optimal utilization of error estimates of data values, Environmetrics, 5(2), 111–126,
- 895 doi:10.1002/env.3170050203, 1994.
- 896 Paatero, P., Eberly, S., Brown, S. G. and Norris, G. A.: Methods for estimating uncertainty in
- 897 factor analytic solutions, Atmos. Meas. Tech., 7(3), 781–797, doi:10.5194/amt-7-781-2014,
 898 2014.
- 899 Parworth, C., Fast, J., Mei, F., Shippert, T., Sivaraman, C., Tilp, A., Watson, T. and Zhang, Q.:
- 900 Long-term measurements of submicrometer aerosol chemistry at the Southern Great Plains
- 901 (SGP) using an Aerosol Chemical Speciation Monitor (ACSM), Atmos. Environ., 106, 43–55,
- 902 doi:10.1016/j.atmosenv.2015.01.060, 2015.
- 903 Petit, J.-E., Favez, O., Sciare, J., Canonaco, F., Croteau, P., Močnik, G., Jayne, J., Worsnop, D. 904 and Leoz-Garziandia, E.: Submicron aerosol source apportionment of wintertime pollution in 905 Paris, France by double positive matrix factorization 906 (PMF<sup&gt;2&lt;/sup&gt;) using an aerosol chemical speciation 907 monitor (ACSM) and a multi-wavelength Aethalometer, Atmos. Chem. Phys., 14(24), 13773– 908 13787, doi:10.5194/acp-14-13773-2014, 2014.
- Pfaffenberger, L., Barmet, P., Slowik, J. G., Praplan, A. P., Dommen, J., Prévôt, A. S. H. and
 Baltensperger, U.: The link between organic aerosol mass loading and degree of oxygenation:
 an α-pinene photooxidation study, Atmos. Chem. Phys., 13(13), 6493–6506, doi:10.5194/acp13-6493-2013, 2013.
- Pope, C. A. and Dockery, D. W.: Health Effects of Fine Particulate Air Pollution: Lines that
 Connect, J. Air Waste Manage. Assoc., 56(6), 709–742,





- 915 doi:10.1080/10473289.2006.10464485, 2006.
- 916 Ramanathan, V., Chung, C., Kim, D., Bettge, T., Buja, L., Kiehl, J. T., Washington, W. M., Fu,
- 917 Q., Sikka, D. R. and Wild, M.: Atmospheric brown clouds: Impacts on South Asian climate
- 918 and hydrological cycle, Proc. Natl. Acad. Sci., 102(15), 5326–5333,
 919 doi:10.1073/pnas.0500656102, 2005.
- Reyes-Villegas, E., Green, D. C., Priestman, M., Canonaco, F., Coe, H., Prévôt, A. S. H. and
 Allan, J. D.: Organic Aerosol source apportionment in London 2013 with ME-2: exploring the
 solution space with annual and seasonal analysis, Atmos. Chem. Phys. Discuss., 1–18,
- 923 doi:10.5194/acp-2016-465, 2016.
- 924 Ripoll, A., Minguillón, M. C., Pey, J., Jimenez, J. L., Day, D. A., Sosedova, Y., Canonaco, F.,
- 925 Prévôt, A. S. H., Querol, X. and Alastuey, A.: Long-term real-time chemical characterization
- 926 of submicron aerosols at Montsec (southern Pyrenees, 1570 m a.s.l.), Atmos. Chem. Phys.,
- 927 15(6), 2935–2951, doi:10.5194/acp-15-2935-2015, 2015.
- 928 von Schneidemesser, E., Monks, P. S., Allan, J. D., Bruhwiler, L., Forster, P., Fowler, D., Lauer,
- 929 A., Morgan, W. T., Paasonen, P., Righi, M., Sindelarova, K. and Sutton, M. A.: Chemistry and
- 930 the Linkages between Air Quality and Climate Change, Chem. Rev., 115(10), 3856–3897,
- 931 doi:10.1021/acs.chemrev.5b00089, 2015.
- 932 Schurman, M. I., Lee, T., Sun, Y., Schichtel, B. A., Kreidenweis, S. M. and Collett Jr., J. L.:
- Investigating types and sources of organic aerosol in Rocky Mountain National Park using
 aerosol mass spectrometry, Atmos. Chem. Phys., 15(2), 737–752, doi:10.5194/acp-15-7372015, 2015.
- Schwarz, J. P., Gao, R. S., Perring, A. E., Spackman, J. R. and Fahey, D. W.: Black carbon
 aerosol size in snow, Sci. Rep., 3, 1–5, doi:10.1038/srep01356, 2013.





- 938 Sug Park, E., Henry, R. C. and Spiegelman, C. H.: Estimating the number of factors to include
- 939 in a high-dimensional multivariate bilinear model, Commun. Stat. Simul. Comput., 29(3),
- 940 723–746, doi:10.1080/03610910008813637, 2000.
- Szidat, S., Prévôt, A. S. H., Sandradewi, J., Alfarra, M. R., Synal, H.-A., Wacker, L. and
 Baltensperger, U.: Dominant impact of residential wood burning on particulate matter in
- Alpine valleys during winter, Geophys. Res. Lett., 34(5), doi:10.1029/2006GL028325, 2007.
- 944The Swiss Federal Council: Ordinance of 16 December 1985 on Air Pollution Control (OAPC).945[online]Availablefrom:https://www.admin.ch/opc/en/classified-
- compilation/19850321/index.html#app7 (Accessed 10 September 2019), 2018.
- 947 Tobler, A., Bhattu, D., Canonaco, F., Lalchandani, V., Shukla, A., Thamban, N. M., Mishra,
- 948 S., Srivastava, A. K., Bisht, D. S., Tiwari, S., Singh, S., Močnik, G., Baltensperger, U., Tripathi,
- 949 S. N., Slowik, J. G. and Prévôt, A. S. H.: Chemical characterization of PM2.5 and source
- 950 apportionment of organic aerosol in New Delhi, India, Sci. Total Environ., 745, 1-12,
- 951 doi:10.1016/j.scitotenv.2020.140924, 2020.
- Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R. and Jimenez, J. L.:
 Interpretation of organic components from Positive Matrix Factorization of aerosol mass
 spectrometric data, Atmos. Chem. Phys., 9(9), 2891–2918, doi:10.5194/acp-9-2891-2009,
 2009.
- Vlachou, A., Daellenbach, K. R., Bozzetti, C., Chazeau, B., Salazar, G. A., Szidat, S., Jaffrezo,
 J.-L., Hueglin, C., Baltensperger, U., Haddad, I. El and Prévôt, A. S. H.: Advanced source
 apportionment of carbonaceous aerosols by coupling offline AMS and radiocarbon sizesegregated measurements over a nearly 2-year period, Atmos. Chem. Phys., 18(9), 6187–6206,
 doi:10.5194/acp-18-6187-2018, 2018.





- 961 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I., Alfarra, M. R.,
- 962 Takami, A., Middlebrook, A. M., Sun, Y. L., Dzepina, K., Dunlea, E., Docherty, K., DeCarlo,
- 963 P. F., Salcedo, D., Onasch, T., Jayne, J. T., Miyoshi, T., Shimono, A., Hatakeyama, S.,
- 964 Takegawa, N., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian,
- 965 K., Williams, P., Bower, K., Bahreini, R., Cottrell, L., Griffin, R. J., Rautiainen, J., Sun, J. Y.,

966 Zhang, Y. M. and Worsnop, D. R.: Ubiquity and dominance of oxygenated species in organic

- 967 aerosols in anthropogenically-influenced Northern Hemisphere midlatitudes, Geophys. Res.
- 968 Lett., 34(13), n/a-n/a, doi:10.1029/2007GL029979, 2007.
- 969 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Ulbrich, I. M., Ng, N. L., Worsnop, D. R. and
- Sun, Y.: Understanding atmospheric organic aerosols via factor analysis of aerosol mass
 spectrometry: a review, Anal. Bioanal. Chem., 401(10), 3045–3067, doi:10.1007/s00216-0115355-y, 2011.
- Zhang, Y., Favez, O., Petit, J.-E., Canonaco, F., Truong, F., Bonnaire, N., Crenn, V., Amodeo,
 T., Prévôt, A. S. H., Sciare, J., Gros, V. and Albinet, A.: Six-year source apportionment of
 submicron organic aerosols from near-continuous highly time-resolved measurements at
 SIRTA (Paris area, France), Atmos. Chem. Phys., 19(23), 14755–14776, doi:10.5194/acp-1914755-2019, 2019.
- Zotter, P., Herich, H., Gysel, M., El-Haddad, I., Zhang, Y., Močnik, G., Hüglin, C.,
 Baltensperger, U., Szidat, S. and Prévôt, A. S. H.: Evaluation of the absorption Ångström
 exponents for traffic and wood burning in the Aethalometer-based source apportionment using
 radiocarbon measurements of ambient aerosol, Atmos. Chem. Phys., 17(6), 4229–4249,
 doi:10.5194/acp-17-4229-2017, 2017.

983







Fig. 1 Chemical composition of PM₁ in Magadino 2013-2014 – daily (a), seasonal (b) and
annual (c) averages. The labels indicate the non-refractory organics (Org), sulphate (SO₄),
nitrate (NO₃), ammonium (NH₄) and chloride (Cl) ions measured by ACSM, and the black
carbon (BC) measured by light absorption.



991 Fig. 2 Seasonal diurnal cycles of PM1 constituents calculated as an hourly average for ACSM











994 Fig. 3 Annual cycle of OA sources: (a) absolute and (b) relative OA contributions plotted as

995 30-min resolved time series, (c) BC source apportionment.







997 Fig. 4 OA pie charts for the whole year and for different seasons.







Fig. 5 Overview of the primary and secondary OA sources in Magadino in 2013-2014: (a) OA factor profiles and (b) seasonal diurnal cycles of HOA, BBOA, LOA, MO-OOA, and LO-OOA. The ambient temperature is shown on the LO-OOA diurnal plots, respectively. In (a) the error bar is the standard deviation; the black bars show the maximum and the minimum that the variable allowed to be vary from the reference profiles. The average, 10th and 90th percentiles for a-values of HOA are, 0.195, 0.007 and 0.378, respectively. Also, the average, 10th and 90th percentiles for a-values of BBOA are 0.202, 0.025 and 0.379, respectively.







1007 **Fig. 6** Diurnal cycles of HOA (grey symbols), black carbon apportioned to traffic emissions 1008 eBC_{tr} (dashed lines) and NO_x (dotted lines) for weekdays (a) and weekends (b). The shaded 1009 areas represent interquartile range for 1-hour average HOA.



1011 **Fig. 7** OOA f_{44} and f_{43} for four different seasons. The yellow cloud of data points represents the 1012 f_{44} vs f_{43} by subtracting the f_{44} and f_{43} contributed from HOA, BBOA and LOA factors. They 1013 are color coded by the total OA mass concentration. The circles, triangles, and squares





- 1014 represent the ratio between f_{44} and f_{43} intensities within the factor profiles of MO-OOA and LO-
- 1015 OOA, respectively. While the smaller size of circles, triangles, and squares are from rolling
- 1016 PMF analysis, which are color coded by the date and time. The dash line are the Sally's triangle
- 1017 from (Ng et al., 2011) and depicts the region where several PMF OOA from the last decade
- 1018 resided in the f_{44} vs f_{43} space.







Fig. 8 Absolute statistical uncertainties of PMF for HOA, BBOA, LOA, LO-OOA, MO-OOA
and total OOA (LO-OOA+MO-OOA) for all data. The data points colour-coded all data points





- 1022 by temperature. The PMF error (uncertainties) of selected PMF runs and rotational 1023 uncertainties is estimated using the slope of the linear regression of standard deviation (σ) vs.
- 1024 the averaged mass concentration (μ) for each factor.
- 1025

1026



Fig. 9 (a) Time series of total oxygenated organic aerosol (LO-OOA+MO-OOA) from online
and offline source apportionment solutions, together with f60 in LO-OOA for online solution,
and levoglucosan in PM10 filter; (b) Averaged LO-OOA factor profile from online solution
during DJF (Dec, Jan, and Feb), when online total OOA is significantly higher than that of
offline solutions.