



1 Vehicle induced turbulence and atmospheric pollution

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7 **Abstract.** Theoretical models of the Earth’s atmosphere adhere to an underlying concept of flow driven by radiative transfer
8 and the nature of the surface over which the flow is taking place: heat from the sun and/or anthropogenic sources are the sole
9 sources of energy driving atmospheric constituent transport. However, another source of energy is prevalent in the human
10 environment at the very local scale – the transfer of kinetic energy from moving vehicles to the atmosphere. We show that
11 this source of energy, due to being co-located with combustion emissions, can influence their vertical distribution to the extent
12 of having a significant influence on lower troposphere pollutant concentrations throughout North America. The effect of
13 vehicle-induced turbulence on freshly emitted chemicals remains notable even when taking into account more complex urban
14 radiative transfer-driven turbulence theories at high resolution. We have designed a parameterization to account for the at-
15 source vertical transport of freshly emitted pollutants from mobile emissions resulting from vehicle-induced turbulence, in
16 analogy to sub-grid-scale parameterizations for plume rise emissions from large stacks. This parameterization allows vehicle-
17 induced turbulence to be represented at the scales inherent 3D chemical transport models, allowing its impact over large regions
18 to be represented, without the need for the computational resources and much higher resolution of large eddy simulation
19 models. Including this sub-grid-scale parameterization for the vertical transport of emitted pollutants due to vehicle-induced
20 turbulence into a 3D chemical transport model of the atmosphere reduces pre-existing North American nitrogen dioxide biases
21 by a factor of eight, and improves most model performance scores for nitrogen dioxide, particulate matter and ozone (for
22 example, reductions in root mean square errors of 20, 9 and 0.5 percent, respectively).

23 1 Introduction

24 A common and ongoing problem with theoretical descriptions of the Earth’s atmosphere is a dichotomy in the representation
25 of turbulent transport, between the turbulence estimated in weather forecast models, and the turbulence required for accurate
26 simulations in air-quality forecast models. Representations of atmospheric turbulence used in weather forecast and climate
27 models have focused on parameterizations of “sub-gridscale turbulence”; descriptions of the storage and release of energy
28 derived from incoming solar radiation and anthropogenic heat release, physical factors in the built-environment, and the
29 transfer of sensible and latent heat between the built environment and the atmosphere. These efforts adhere to an underlying
30 concept of radiative-driven flow: heat transfer from the sun and/or anthropogenic sources being the source of energy behind



31 atmospheric motions. There has been considerable research focused on improving understanding radiative-driven flow in
32 urban areas (e.g. the advection and diffusion associated with buildings and street canyons (Mensink *et al.*, 2014), urban heat
33 island radiative transfer theory (Mason *et al.*, 2000), and in efforts to increase 3D model vertical and horizontal resolution in
34 order to better capture the physical environment (Leroyer *et al.*, 2014). However, when these physical models of turbulence
35 are applied to problems involving the emissions, transport and chemistry of atmospheric pollutants, predicted surface
36 concentrations of emitted pollutants may be biased high, and concentrations aloft biased low, indicating the presence of missing
37 additional sources of atmospheric dispersion (Makar *et al.*, 2014; Kim *et al.*, 2015). Despite ongoing work to improve the
38 turbulence schemes in meteorological models (Makar *et al.*, 2014; Hu *et al.*, 2013; Klein *et al.*, 2014), computational predictive
39 models of atmospheric pollution typically make use of a constant “floor” or “cut-off” in the thermal turbulent transfer
40 coefficients provided by weather forecast models, sometimes with higher values of this cutoff over urban compared to rural
41 areas (Makar *et al.*, 2014), in an attempt to compensate for apparent insufficient vertical mixing of chemical tracers. The
42 turbulent mixing in these physical descriptions, while capable of reproducing observed meteorological conditions, do not
43 explain lower concentration observations of emitted atmospheric pollutants.

44 Large stack sources of pollutants provide a useful analogy in investigating the cause of this discrepancy. Emissions from these
45 sources occur at high temperatures, lofting their emitted mass high into the atmosphere as a result of buoyancy effects.
46 However, the physical size of the stacks (< 10 m diameter) is much smaller than the grid cell size used in regional models (km
47 to 10’s of km). In order to capture the rapid vertical redistribution of emissions from large stacks, sub-grid-scale
48 parameterizations are used, in which buoyancy calculations are performed to determine plume heights, which are then used to
49 determine the distribution of freshly emitted pollutants (Briggs, 1975; Briggs, 1984; Gordon *et al.*, 2018; Akingunola *et al.*,
50 2018). For large stack emissions, these parameterizations account for the effect of the addition of energy (the hot exhaust gas)
51 on the local distribution of pollutants, and are essential in estimating initial vertical distribution of those pollutants.

52 In this work, we investigate the potential for another type of at-source energy to influence the vertical distribution of freshly
53 emitted pollutant concentrations: the addition of kinetic energy due to the displacement of air during the passage of vehicles
54 on roadways. Roadway observations in the 1970’s showed that this transferred energy has a significant influence on the
55 transport of primary pollutants released from vehicle exhaust, with vehicle passage being associated with “a distinct bulge in
56 the high frequency range of the wind spectrum”, “corresponding to eddy sizes on the order of a few metres” (Rao *et al.*, 1979).
57 The same work found that the variation in the concentration of non-reactive tracers could be attributed to wakes behind moving
58 vehicles. Subsequent theoretical development led to the creation of the roadway-scale models describing turbulence within a
59 few 10’s of metres around and above roadways, in turn used to estimate the very local-level impact of vehicles on emitted
60 pollutant concentrations (Eskridge and Catalano, 1987). These models showed that near-roadway concentrations of emitted
61 pollutants were highly dependent on vehicle speed, with over a factor of two reduction in emission-normalized pollutant
62 concentrations being associated with an increase in vehicle speed from 20 to 100 km/hr (Eskridge *et al.*, 1991). With the
63 advent of portable, very high time resolution 3-D sonic anemometers, the turbulent kinetic energy of individual vehicles could
64 be measured directly, either aboard an instrumented trailer towed behind a vehicle¹¹, or through instrumentation mounted



65 aboard a laboratory following other vehicles in traffic (Gordon *et al.*, 2012; Miller *et al.*, 2018). However, the application of
66 this information has been limited up to now to theoretical and computational models of the near-roadway environment.
67 Regional air-quality models also have vertical resolution in the 10's of metres near the surface, suggesting the potential for
68 vehicle-induced turbulence (VIT) to influence turbulent mixing out of the lowest model layer(s). Here we demonstrate that
69 this sub-grid-scale vertical transport process, which due to its highly localized spatial nature (over roadways), has a
70 disproportionate impact on the vertical distribution and transport of freshly emitted chemical tracers. A comparable sub-grid-
71 scale process which has a similar influence on pollutants are the emissions from large stacks (Gordon *et al.*, 2018; Akingunola
72 *et al.*, 2018). Accurate estimation of pollutant concentrations from the latter sources must take into account the at-source
73 buoyancy and exit velocity of high-temperature exhaust to determine the vertical distribution of fresh emissions. Similarly,
74 our work focusses on determining the local lofting of pollutants from and due to moving vehicles, in order to adequately
75 represent the at-source vertical distribution of their emissions, on the larger scale.

76 The extent of the vertical influence of VIT varies depending on the configuration of vehicles on the roadway. From
77 observations taken from a trailer following an isolated passenger van (Rao *et al.*, 2002), and large eddy simulation (LES) /
78 computational fluid dynamics (CFD) models of individual vehicles (Kim *et al.*, 2011; Kim *et al.*, 2016), the vertical distance
79 over which VIT can be distinguished from the background for isolated, *individual* vehicles (i.e. the mixing length) is on the
80 order of 2.5 to 5.13 m. However, as we show in Methods and Results, for observations of *ensembles* of vehicles in traffic
81 (Gordon *et al.*, 2012; Miller *et al.*, 2018), and large eddy / computational fluid dynamics simulations of *ensembles* of vehicles
82 (Kim *et al.*, 2016; Woodward *et al.*, 2019; Zhang *et al.*, 2017), the mixing lengths associated with VIT are larger, on the order
83 of 10's of m, to as much as 41 m. The vertical extent of the impacts of alternating low and high areas of surface roughness
84 have been shown to create downwind internal boundary layers to even more significant heights in the atmosphere (e.g 300m,
85 Bou-Zeid *et al.*, 2004, Figure 12), suggesting that impacts into the lower boundary layer due to the alternating roughness
86 elements (in our case, vehicles versus roadways) is not unreasonable. We also show in Methods that the impact of VIT within
87 the context of an air-quality model is via changes to the vertical gradient of the thermal turbulent transfer coefficients; the
88 gradient of the sum of the natural turbulence and VIT terms, allows VIT to influence vertical mixing, even when model vertical
89 resolution is relatively coarse.

90 Large eddy simulation (LES) / computational fluid dynamics (CFD) models have shown the importance of VIT towards
91 modifying local values of turbulent kinetic energy, as noted in the references above. However, these models require very small
92 grid cell sizes and time steps to allow forward time stepping predictions of future meteorology and chemistry. These
93 constraints in turn severely limit the size of the domain in which they can be applied, and the processing time for simulations
94 for these reduced domains can be very high. For example, the FLUENT model was used by Kim *et al.* (2016) with an adaptive
95 mesh with a minimum cell size of 1 cm, with a 100x20x20m domain, while Woodward *et al.* (2019)'s implementation of
96 FLUENT had an equivalent cell size of 50 cm, operating in a domain of 600,000 nodes (a volume of 75,000 cubic metres),
97 and an adaptive timestep limited by a Courant number of 5. The latter criteria implies a computation timestep of less than 0.09
98 s for a 100 km hr⁻¹ vehicle (or wind) speed, while a 1 cm grid cell size implies a computation timestep of less than 1.8x10⁻³ s



99 timestep. Similarly, the LES model employed by Zhang et al (2017) utilized a 1m x 2m x 1m cell size and a computation
100 timestep of 0.03 s. In contrast, a 3D regional chemical transport model typically operates over a domain with may be
101 continental in extent, with limiting horizontal resolution on the order of kilometres, a limiting vertical resolution on the order
102 of 10's of metres, and timesteps on the order of 1 minute. These limits for regional chemical transport models are a function
103 of the need to provide chemical forecasts over a relatively large region, within a reasonable amount of current supercomputer
104 processing time. However, a “scale gap” exists between LES and regional chemical transport models – for regional chemical
105 transport models, parameterizations of the physical processes resolvable at the very high resolution of LES models are required.
106 In return, these parameterizations allow the relative impact of the parameterized processes on the urban to regional to
107 continental scales regional chemical transport models to be determined.

108 Here we make use of both the observational and LES modelling studies to devise a parameterization for VIT, which we then
109 apply at two different configurations of a regional chemical transport model (GEM-MACH). We show that VIT has a
110 potentially significant impact on pollutant concentrations at the urban, regional, and continental scales. Reductions of model
111 biases are of particular interest from the standpoint of the use of air-quality model predictions to determine chronic health
112 outcomes. The inclusion of VIT reduces the positive bias in predictions of North American and urban scale nitrogen dioxide
113 by factors of 8 and 2.6 respectively, and improves the accuracy of model simulations for most statistics for nitrogen dioxide,
114 ozone and particulate matter.

115 **2 Methodology**

116 **2.1 Theoretical development**

117 In contrast to the very local resolution “roadway” models used to examine the impact of vehicle motion on pollutant
118 concentration (Eskridge and Catalano, 1987; Eskridge *et al.*, 1991), and computational fluid dynamics modelling of vehicle
119 turbulence (Kim *et al.*, 2011; Kim *et al.*, 2016; Woodward *et al.*, 2019; Zhang *et al.*, 2017), 3D models of atmospheric pollution
120 (Galmarini *et al.*, 2015) have horizontal grid-cell sizes of a one to 10's of km, and thus emissions and vertical transport
121 associated with roadways must be approached from the standpoint of sub-grid-scale parameterizations. Measurements of the
122 turbulent kinetic energy (TKE) associated with vehicles are usually available on a “per-vehicle” or “per-vehicle within an
123 ensemble” basis. These observations provide the average on-road TKE per vehicle passing a point per unit time (Gordon *et al.*,
124 2012; Miller *et al.*, 2018) and/or the shape of the enhanced TKE cross-section in the plane perpendicular to the vehicle's
125 motion (Rao *et al.*, 2002). A sub-gridscale parameterization linking these scales is therefore necessary in order to study the
126 impacts of VIT on the vertical redistribution of freshly emitted pollutants, and hence on large-scale atmospheric chemistry and
127 transport. Sub-gridscale parameterizations are commonly used in atmospheric models of weather forecasting to provide the
128 rates of change of processes which occur at scales smaller than the model's horizontal and/or vertical resolution: cloud
129 formation and buoyant plume rise from large stacks being a common example for model grid cell sizes of 10km or more (Kain,
130 2004; Briggs, 1975; Briggs 1984; Gordon *et al.*, 2018; Akingunola *et al.*, 2018).



131 Three separate problems must be addressed in the construction of such a VIT parameterization for atmospheric chemical
132 transport models, specifically:

- 133 (1) What is the relationship governing the decrease in VIT with increasing distance (height) from the vehicles?
134 (2) How can observation data, in units of vehicles per unit time, be related to variables more commonly available for
135 regional chemical transport models?
136 (3) How can VIT be incorporated into a regional model in a manner that only the emissions due to vehicles are affected,
137 given that the vehicle-induced turbulence will have the most significant impact on emissions from moving vehicles due to the
138 relatively low area fraction of roadway area within a given grid cell?
139 We address each of these issues in the sub-sections that follow.

140 2.2 Changes in VIT with Height

141 Measurements of TKE behind a passenger van (Rao *et al.*, 2002) typically show a smooth distribution, with TKE decreasing
142 both above and below the height of the upper trailing edge of the moving vehicle. Similar results have been seen from very
143 high resolution computational fluid dynamics modelling of the flow around individual vehicles, though the shape of the vehicle
144 and the arrangement of vehicles on the roadway can have a strong influence on the location of the maximum and shape of the
145 vertical profile in TKE (Kim *et al.*, 2011; Kim *et al.*, 2016). We examined four datasets (Rao *et al.*, 2002; Kim *et al.*, 2016;
146 Woodward *et al.*, 2019; Zhang *et al.*, 2017) to evaluate the extent to which a Gaussian distribution may be used to represent
147 the decrease in VIT with height above moving vehicles, as well as examining the expected range of mixing lengths which may
148 result from VIT. A Gaussian distribution of TKE with height is given by equation (1), where $I_q(z)$ is the time integrated added
149 TKE value for vehicle type q with height z (m^2s^{-1}), h_q is the height of the vehicle, and A_q and σ_q are numerical constants:

$$150 \quad I_q(z) = \frac{A_q}{\sqrt{2\pi\sigma_q^2}} e^{\left(-\frac{(z-h_q)^2}{2\sigma_q^2}\right)} \quad (1)$$

151 Equation (1) may be re-written as:

$$152 \quad \ln(\sqrt{2\pi}I_q(z)) = \ln\left(\frac{A_q}{\sigma_q}\right) - \frac{(z-h_q)^2}{2\sigma_q^2} \quad (2)$$

153 Equation (2) shows that values of $-(z-h_q)^2$ versus $\ln(\sqrt{2\pi}I_q(z))$, with the values of z taken from vertical profiles of $I_q(z)$
154 in the literature, will yield a slope of $\frac{1}{2\sigma_q^2}$ and an intercept of $\ln\left(\frac{A_q}{\sigma_q}\right)$, and the correlation coefficient for this relationship may
155 be used to judge the accuracy of the use of a Gaussian distribution to describe the decrease in TKE with height above moving
156 vehicles. The resulting relationships may also be used to describe the vertical mixing length, defined “as the diameter of the
157 masses of fluid moving as a whole in each individual case; or again, as the distance traversed by a mass of this type before it
158 becomes blended in with neighbouring masses” (Prandtl, 1925; Bradshaw, 1974). Here we assume that this blending has



159 occurred at the height at which the Gaussian has dropped to 0.01 of the value at $z=h_q$ (i.e. the value of z at which VIT has

160 reached 1% of its maximum value (i.e. $e^{\left(\frac{(z-h_q)^2}{2\sigma_q^2}\right)} = 0.01$).

161 An example of the analysis used to construct Table 1 appears below (Figure 1), for a CFD example for an ensemble of vehicles,
162 taken from the literature (Kim *et al.*, 2016). In this figure, contours of TKE are shown as solid lines. TKE values as a function
163 of height at three locations behind the trucks were used to determine σ_q and hence estimate the length scale via equations (1)
164 and (2). A notable feature of this example is the substantial increase in length scale which occurs between the initial vehicle
165 (a transport truck) and subsequent downwind vehicles (compare height of TKE contours, and the resulting length scales in
166 Figure 1, between left and right sides of the figure). Increases in downwind turbulent length scales associated with vehicles
167 moving in close ensembles are a common feature in the literature.

168

169 This analysis (see Table 1) shows that a Gaussian distribution accounts for much of the variability in TKE with height
170 (correlation coefficients of 0.54 to 0.99), and under realistic traffic conditions, the mixing lengths increase in size, and may be
171 considerably larger than those of isolated vehicles.

172 Two VIT mobile laboratory studies (Gordon *et al.*, 2012; Miller *et al.*, 2018) observed vehicle-per-second TKE for vehicles
173 moving in ensembles along multilane roadways, aggregated by vehicle classes using the same methodology, to derive formulae
174 for the net TKE added by VIT at 4m and 2m (the height of the instrumentation used in these studies). We combine these data
175 here to determine the change in VIT with height. Setting E as the TKE added due to the vehicles, two formulae result:

176

$$\begin{aligned} E(4m) &= 1.8 F_c + 2.2 F_m + 20.4 F_t \\ E(2m) &= 2.4 F_c + 6.2 F_m + 14.8 F_t \end{aligned} \quad (3)$$

177 Where $E(4m)$ and $E(2m)$ are the TKE added driving within the ensemble at 4 and 2 m elevation from these two studies ($\text{m}^2 \text{s}^{-2}$), and F_c , F_m , and F_t are the number of passenger cars, mid-sized (vans, flatbed pickup trucks, and SUVs) and large vehicles
178 (10 to 18 wheel heavy-duty vehicles) travelling past a given point on the highway per second. The numerical coefficients are
179 the time integrated TKE values (I_q) at the two heights (m^2s^{-1}). An alternative approach would be to make use of vehicle speed
180 data within each grid cell and parameterizations utilizing vehicle speed (Di Sabatino *et al.*, 2003; Kastner-Klein *et al.*, 2003)
181 to construct TKE additions due to the sub-grid-scale roadways. However, vehicle speed information is not currently readily
182 available on a gridded hourly basis, while estimates of vehicle km travelled are available in gridded form due to their use in
183 emissions processing, and making the simple scaling assumption that the vehicles travel across one dimension of a grid cell
184 allows us to generate the F_c values required to estimate TKE. Note that vehicle speed is implicit in this methodology utilizing
185 VKT – higher speeds will result in a greater number of vehicle km travelled per unit time, and hence higher TKE values. As
186 in the above discussion, we assume a Gaussian distribution of the coefficients of the TKE equations of (3) with height for each
187 vehicle, where $h_q = 1.5\text{m}$, 1.9m and 4.11m for cars, mid-sized vehicles and trucks, respectively, with each of the 2m and 4m
188 values of the coefficients of (3) being used to determine the corresponding values of A_q and σ_q of equation (1), (i.e. $q = c, m, t$).
189 The resulting height-dependent formulae may be used to replace the coefficients of (3), leading to the following formula for
190



191 the net turbulent kinetic energy associated with the number of vehicles in transit along a given stretch of roadway at a given
 192 time:

$$\begin{aligned}
 E_{net}(z) = & 2.43F_c e^{[-2.40 \times 10^{-2}(z-1.5)^2]} \\
 & + 15.58F_m e^{[-1.18 \times 10^{-1}(z-1.9)^2]} \\
 & + 20.43F_t e^{[-3.61 \times 10^{-2}(z-4.11)^2]}
 \end{aligned} \quad (4)$$

194 Most 3-D chemical transport models make use of some variation of “K-theory” diffusion to link turbulent kinetic energy to
 195 mixing, with the vertical mixing of a transported variable c due to turbulence at heights z being related to the thermal turbulent
 196 transfer coefficient K via:

$$\frac{\partial c}{\partial t} = \frac{\partial}{\partial z} \left(K \frac{\partial c}{\partial z} \right) \quad (5)$$

198 Finite differences and tridiagonal matrix solvers are usually used to forward integrate equation (5). For example, the solver
 199 used in the GEM-MACH model uses the following finite difference for the spatial derivatives (both spatial derivatives are
 200 $O(\Delta\sigma^2)$, the derivatives are carried out in, and the K values are transformed into, $\sigma = \frac{P}{P_0}$ coordinates as \tilde{K} , where P is the
 201 pressure, and P_0 is the surface pressure):

$$\frac{c_i^{n+1} - c_i^n}{\Delta t} = \frac{\frac{1}{2}(\tilde{K}_{i+1} + \tilde{K}_i) \left(\frac{c_{i+1} - c_i}{\sigma_{i+1} - \sigma_i} \right) - \frac{1}{2}(\tilde{K}_i + \tilde{K}_{i-1}) \left(\frac{c_i - c_{i-1}}{\sigma_i - \sigma_{i-1}} \right)}{\sigma_{i+\frac{1}{2}} - \sigma_{i-\frac{1}{2}}} \quad (6)$$

203 Note in (6) that the prognostic values of K calculated by the weather forecast model are on the same vertical levels as
 204 concentration; values of the additional component of K associated with VIT must therefore be calculated for model layers as
 205 opposed to layer interfaces.

206 K and E may be linked through the relationship of Prandtl, where l is a characteristic length scale:

$$K = 0.4 l \sqrt{E} \quad (7)$$

208 As was done for Table 1, we have chosen this value on a per-vehicle basis as the vertical location at which the Gaussian
 209 profiles derived above reach 0.01 (i.e. 1%) of their maximum value. Using each of the coefficient values of (3) at the two
 210 heights, in conjunction with equation (1) treated as a two-variable in two unknowns (A_q , σ_q) problem we find values of l_c , l_m ,
 211 and l_t of 13.56, 6.25, and 11.28 m, respectively. These values are based on observed traffic conditions, and fall well within
 212 the range of mixing lengths provided for vehicle ensembles in Table 1, however, we note that they are a source of uncertainty,
 213 with the percent uncertainties (Gordon *et al.*, 2012) associated with the 4m values at $\pm 52\%$, $\pm 157\%$, and $\pm 12\%$ for cars, mid-
 214 sized vehicles and trucks, respectively. The relatively low values of l_m and high uncertainties in the corresponding mid-sized
 215 vehicle per-vehicle estimates of TKE relative to the other vehicle types are likely the result of a combination of small sample
 216 size (Gordon *et al.* (2012) noted the relative proportion of the three vehicle classes as 89.9% cars, 4.8% mid-sized, and 5.3%
 217 trucks, respectively) and the variety of ensemble versus isolated vehicles sampled (noting the variation in Table 1 for vehicles
 218 within the smaller vehicle size classes). Additional observations of vehicle turbulence are clearly needed, particularly in the
 219 region above the largest vehicles on the road (4.1m), using remote sensing techniques such as Doppler lidar, in order to improve
 220 mixing length estimates. However, the values used here are reasonable with respect to the available data, and while likely



221 overestimating the mixing length associated with isolated vehicles (Rao *et al.*, 2002; Kim *et al.*, 2016) likely underestimate
222 the mixing length of ensembles of vehicles (Kim *et al.*, 2016), particularly for ensembles moving within street canyons
223 (Woodward *et al.*, 2019; Zhang *et al.*, 2017). The latter represent the some of the specific regions where vehicle emissions are
224 likely to dominate.

225 We derive the following formula for the addition to the thermal turbulent transfer coefficient associated with vehicle passage
226 as a function of height:

$$227 \quad K_{VIT}(z) = 0.4 \frac{l_c F_c + l_m F_m + l_t F_t}{F_c + F_m + F_t} \sqrt{\begin{matrix} 2.43 F_c e^{[-2.40 \times 10^{-2}(z-1.5)^2]} \\ + 15.58 F_m e^{[-1.18 \times 10^{-1}(z-1.9)^2]} \\ + 20.43 F_t e^{[-3.61 \times 10^{-2}(z-4.11)^2]} \end{matrix}} \quad (8)$$

228
229 The use of (8) must be undertaken with care. Like most regional air-quality models, the vertical resolution of GEM-MACH
230 used here is relatively coarse (the first four model layer midpoints are located approximately 24.9, 99.8, 205.0, and 327.0 m
231 above the surface). Layer midpoint values must be representative of the layer resolution in order to describe the impact of
232 VIT on the layer. A simple linear interpolation between the peak values of K_{VIT} and the first model interface will overestimate
233 the impact of VIT within the lowest model layer, while the use of (8) for the mid-point value alone will underestimate the
234 influence of VIT within the lowest part of the first model layer. The best representation of a sub-grid-scale scalar quantity
235 within a discrete model layer is its vertical average within that layer. Here, we calculate the vertically integrated average of
236 (8) within each model layer, to provide the best estimate of the impact of VIT, to within the vertical resolution of the model.

237 2.3 VIT and Model Vertical Resolution

238 The issue of the vertical extent of the impact of VIT is worth considering in the context of model layer thickness. Given that
239 the vertical length scale of added VIT is on the order of 10's of metres, as denoted in the studies quoted herein, it is reasonable
240 to question whether the added turbulence should be expected to have an impact on the dispersion of pollutants. This apparent
241 contradiction is easily resolved by noting, (1) that the turbulence due to VIT is added as an addition to the pre-existing
242 “meteorological” thermal turbulent transfer coefficient (with the *net* turbulence profile, not the VIT alone, determining its
243 impact on vertical mixing); and (2) that the impact of this net turbulence does not depend just on the magnitude of the net
244 coefficients of thermal turbulent transfer, but also on their vertical gradient. This second point can be illustrated by expanding
245 the diffusion equation using the chain rule of calculus (i.e. $\frac{\partial c}{\partial t} = \frac{\partial}{\partial z} \left(K_{net} \frac{\partial c}{\partial z} \right) = K_{net} \frac{\partial^2 c}{\partial z^2} + \frac{\partial K_{net}}{\partial z} \frac{\partial c}{\partial z}$), and the aid of an example,
246 shown in Figure 2. Figure 2 displays examples of cases where the concentration gradient and natural thermal turbulent transfer
247 coefficient both decrease linearly with height (Figure 2(a,b)), and where the concentration gradient decreases with height while
248 the natural thermal turbulent transfer coefficients increase with height (Figure 2(c,d)). The added K_{VIT} is shown as a blue
249 dashed line, and the net vertical thermal turbulent transfer is shown as a red line. Figure 2 (a) and Figure 2(c) depict these
250 curves at a high vertical resolution, while Figure 2(b) and Figure 2(d) depict them at a low (regional model) resolution. Note



251 that in the latter, the vehicle-induced addition to the net thermal turbulent transfer coefficient depicted in Figure 2(a,c) lies
252 entirely within the lowest model layer of Figure 2(b,d). In both Figure 2(a) and Figure 2(b), the impact of K_{VIT} is to slow the
253 build-up of near-surface concentrations. In both Figure 2(c) and Figure 2(d), the impact of K_{VIT} is to more rapidly vent near-
254 surface concentrations further up into the atmosphere. That is, at both high and low resolution, K_{VIT} affects near-surface
255 concentrations, due to the vertical gradient of $\frac{\partial K_{net}}{\partial z}$. Centered difference calculations for the low resolution case are shown
256 in Figure 2(b,d) to illustrate the point that gradients in low vertical resolution net diffusivity result in reductions in lowest
257 model layer trapping, and increases in venting from this lowest layer. In both of these cases, the addition of vehicle turbulence
258 to the lowest model layer changes the gradient of the net thermal turbulent transfer coefficient, in turn leading to reduced
259 surface concentrations. The above example illustrates the manner in which VIT may have an impact even on relatively low
260 vertical model resolution.

261 **2.4 Relating VIT to Available Gridded Data – Vehicle Km Travelled**

262 Along individual roadways, the equation (8) makes use of F_c , F_m , and F_t observations at points along roadways within a grid-
263 cell, hence deriving local estimates of VIT. This data is currently difficult to obtain for large-scale applications, and hence
264 we have turned to secondary sources of information to estimate these three terms. Vehicle Kilometer Travelled (VKT) is used
265 for estimating onroad vehicle emissions at jurisdiction level (e.g. county level for the US and province level for Canada) for
266 the national emissions inventory. Emissions processing systems used for air-quality models make use of spatial surrogates to
267 help determine the spatial allocation of the mass emitted from different types of vehicles on different roadways (Adelman *et*
268 *al.*, 2017). The same set of surrogates is used for calculating VKT (km s^{-1}) for each grid cell of the model domain (varying
269 by hour of day and day of week, for each of the three vehicle categories listed (see Figure 3), in turn providing diurnal variations
270 of VIT matching traffic flow. The data shown are derived from 2006 Canadian (Taylor, 2019) and 2011-based projected 2017
271 US VKT (EPA, 2017). Note that for the 10km grid cell size used here, values of F_c , F_m , and F_t may be derived by dividing
272 these numbers by 10. The largest contribution to total vehicle km travelled is by the “cars” class (Figure 3(a)) due to their
273 greater numbers (the originating study (Miller *et al.*, 2018) found that 89.9% of vehicles measured were cars), followed by
274 trucks (Figure 3(c); 5.3% of vehicles measured), and mid-sized vehicles (Figure 3(b); 4.8% of vehicles measured).
275 These VKT data may be linked to the above VIT formula (8), provided the distance each vehicle is travelling within that grid
276 cell is known. Here, we have made two additional assumptions. The first assumption is that each vehicle carries out a simple
277 transit of the cell – the distance travelled is the cell-size. While this may be a reasonable first-order approximation, we note
278 that it has limitations: for example, when the number of vehicles on the roads overwhelm the capacity of the roads (rush-hour
279 traffic jams) the distance travelled decreases. However, under these circumstances the VKT values will also decrease; the
280 impact of rush-hour conditions should to some extent be included within the VKT estimates available for emissions processing
281 systems. The second assumption is that the VKT contributions within a grid-cell are additive – i.e. that their numbers may be
282 added via the “F” terms in (6) (Gordon *et al.*, 2012; Miller *et al.*, 2018), an assumption found to be accurate in CFD modelling



283 (Kim *et al.*, 2016). Note that this assumption may result in overestimates of the net TKE – a better methodology for future
284 work would be to collect and make use of statistics of vehicle density by roadway type within each grid-cell. However, we
285 note that assuming that vehicles are evenly distributed over roadways in a grid cell would result in a net underestimate of the
286 TKE contributed over the larger roadways and main arteries of urban areas.

287 Example 10 AM EDT North American 10km resolution gridded vehicle-induced thermal turbulent transfer coefficient values
288 (K_{VIT} , equation 8) created using these assumptions, and an example vertical profile of K_{VIT} for central Manhattan Island at
289 0.5m vertical resolution are shown in Figure 4. The resulting enhancements to “natural” K values at the vertical resolution of
290 the version of the GEM-MACH air-quality model, at 2.5km horizontal resolution, are shown in Figure S1 as dashed lines. The
291 enhancements are confined to the lowest model layer, as might be expected from the vertical resolution employed in this
292 version of GEM-MACH. Nevertheless, the values are sufficient to significantly change simulated vertical transport due to
293 modifications to the resolved gradient in thermal turbulent transfer coefficients, as discussed above. Both the magnitude and
294 gradient of $K_{net} = K + K_{VIT}$ may contribute to the concentration changes: breaking the vertical diffusion equation down using
295 the chain rule, (5) may be rewritten

$$296 \quad \frac{\partial c}{\partial t} = K \frac{\partial^2 c}{\partial z^2} + \frac{\partial K}{\partial z} \frac{\partial c}{\partial z} \quad (9)$$

297 Both terms on the right-hand-side of (9) may contribute to decreases in concentration c at the surface and increases in
298 concentrations aloft. If the near-surface concentration profile ($\partial c / \partial z$) is negative (concentrations decrease with height), then
299 increases in K will result in surface concentration decreases). If this results in sufficient lofting that the concentration profile
300 maximizes above the ground (i.e. $\partial c / \partial z$ becomes positive near the surface), then decreasing values of K with height (i.e.
301 negative values of $\partial K / \partial z$) will also result in a shift towards negative rates of change, through the second term in the right-hand-
302 side of (9). All six panels of Figure S1 show increased K values; i.e. increases in the first term in (9). All six panels also
303 show a trend of $\partial K / \partial z$ becoming more negative, decreasing the magnitude of the second term in (9) in Figure S1 (b,c,d,f), and
304 switching to a negative rate of change in Figure S1(a,e). Both changes in the magnitude and gradient of K resulting from VIT
305 contribute to the resulting changes in surface concentration.

306 The thermal turbulent transfer coefficient values of Figure S1 may also be compared to the minima on “natural” K values
307 imposed in air pollution models in an attempt to account for missing subgrid-scale mixing (Makar *et al.*, 2014; these are
308 typically on the order of 0.1 to 2.0 m^2s^{-1}). Aside from Figure S1(a), the vertical profiles here would not be modified by these
309 lower limits. We also note that these VIT-induced changes in total thermal turbulent transfer coefficients only impact the
310 species emitted at the road-way level, as discussed below.

311 2.5 Construction of a Sub-Gridscale Parameterization for On-Road Vehicle-Induced Turbulence

312 We note that the portion of the area of a grid-cell which is roadway-covered will be relatively small for most air pollution
313 model resolutions, such as those considered here. For example, satellite imagery of the largest freeways show these to have
314 a width of less than 400m. Hence, the largest roads make up less than 1/5 of the total area of a 2.5km grid-cell, and less than



315 1/20 of a 10km grid cell). The largest impact of VIT is thus likely to be for the chemical species being emitted by the mobile
316 sources, in terms of the grid-cell average concentration. Furthermore, the grid cell approach common to these models results
317 in horizontal numerical diffusion from the roadway scale to the grid cell scale: sub-grid-cell scale emissions are automatically
318 mixed across the extent of the grid cell. The key impact of VIT will thus be in the vertical dispersion of the pollutants emitted
319 from mobile sources. We must therefore devise a numerical means to ensure this additional source of diffusion is added to the
320 model, bearing these constraints in mind.

321 Two examples of similar sub-gridscale processes appear in the literature. The first example are the cloud convection
322 parameterizations used in numerical weather forecast models (Kain *et al.*, 2004), wherein the formation and vertical transport
323 associated with convective clouds, are known to occur at smaller scales than the grid cell size employed in a numerical weather
324 prediction model, are treated using sub-gridscale parameterizations. In these parameterizations, cloud formation and transport
325 are calculated within the grid-cell on a statistical basis, using formulae linking the local processes to the resolvable scale of the
326 model. The second example is found in the treatment of emissions from large stacks within air-quality forecast models (Gordon
327 *et al.*, 2018; Akingunola *et al.*, 2018). These sources usually have stack diameters on the order less than 10m, and these sources
328 emit large amounts of pollutant mass at high temperatures and velocities. In order to represent these sources, the most common
329 approach is to calculate the height of the buoyant plume using the predicted ambient meteorology (vertical temperature profile,
330 etc.) as well as the stack parameters (exit velocity, exit temperature, stack diameter). The emitted mass during the model
331 timestep from the stack is then distributed over a defined vertical region within the gridcell in which the source resides. Note
332 that the mass is also automatically distributed immediately in the horizontal dimension within the grid cell – the key issue is
333 to ensure that the emitted mass is properly distributed in the vertical dimension. Our aim in the VIT parameterization that
334 follows is identical in intent to that of the existing major point source treatments in air-quality models: to redistribute the mass
335 emitted by vehicle sources in the vertical dimension, taking the very local physics influencing that vertical transport of fresh
336 emissions into account. We therefore focus on determining the at-source vertical transport of emitted mass associated with
337 VIT.

338 We start with the formulae for the transport of chemical species by vertical diffusion:

$$339 \quad \frac{\partial c_i}{\partial t} = \frac{\partial}{\partial z} \left(K \frac{\partial c_i}{\partial z} \right) + E_i \quad (10)$$

340 Where c_i is the emitted chemical species, K represents the sum of all forms of thermal turbulent transfer in the grid-cell, and
341 E_i is the emissions source term for the species emitted at the surface (applied as a lower boundary condition on the diffusion
342 equation). For grid-cells containing roadways and hence mobile emissions, we split K into meteorological and vehicle-induced
343 components (K_T and K_{VIT} respectively), and the emissions into those from mobile sources and those from all other sources
344 ($E_{i,mob}$ and $E_{i,oth}$, respectively):

$$345 \quad \frac{\partial c_i}{\partial t} = \frac{\partial}{\partial z} \left[(K_T + K_{VIT}) \frac{\partial c_i}{\partial z} \right] + E_{i,mob} + E_{i,oth} \quad (11)$$

346 The terms in (11) may be rearranged:

$$347 \quad \frac{\partial c_i}{\partial t} = \left\{ \frac{\partial}{\partial z} \left[K_T \frac{\partial c_i}{\partial z} \right] + E_{i,oth} \right\} + \left\{ \frac{\partial}{\partial z} \left[(K_T + K_{VIT}) \frac{\partial c_i}{\partial z} \right] + E_{i,mob} \right\} - \left\{ \frac{\partial}{\partial z} \left[K_T \frac{\partial c_i}{\partial z} \right] \right\} \quad (12)$$



348

349 The first bracketed term in (12) describes the rate of change of the chemical due to its emission by non-mobile area sources
350 and vertical diffusion due to meteorological sources of turbulence within the grid-cell, but outside of the sub-grid-scale
351 roadway. The second term describes the rate of change of the vertical diffusion of the mobile-source-emitted pollutants over
352 the sub-grid-cell roadway, which experiences both meteorological and roadway turbulence, and the final term prevents double-
353 counting of the meteorological component in equation (11), which is equivalent to equation (12). Note that turbulent mixing
354 for non-emitted chemicals is determined by solving equation (5), and for chemicals which are not emitted from mobile on-
355 road sources, equation (10) is solved, with $E_i = E_{i,oth}$. This form of the diffusion equation (12) allows the net change in
356 concentration to be calculated from three successive calls of the diffusion solver, starting from the same initial concentration
357 field. One advantage of this approach is that existing code modules for the solution of the vertical diffusion equation may be
358 used – rather than being used once, they are used three times, with different values for the input coefficients of thermal turbulent
359 transfer coefficient (K). The solution, once a suitable means of estimating K_{VIT} is available, is thus relatively easy to implement
360 in existing numerical air pollution model frameworks.

361 2.6 Comparison of energy densities: VIT, Solar, and Urban Perturbations in Sensible and Latent Heat

362 The relative contribution of TKE from VIT towards energy density can be compared to the daytime solar maximum energy
363 input to illustrate why VKT has relatively little impact during daylight hours, particularly in the summer. The maximum TKE
364 from VIT can be determined easily from Figure 3 and the use of our formulae; Figure 3(a) shows vehicle km travelled values
365 ranging from a maximum of 308 in the highest density 10km grid cell in North America (New York City) down through four
366 orders of magnitude in background grid cells with few vehicles. A typical urban value would be 30.8 VKT: this gives an F_c
367 value from our formulae of 3.08 vehicles s^{-1} for a 10km grid cell size. Assuming that the vehicles are all cars, from our
368 formulae we have a corresponding total TKE added at the point crossed by the vehicles, at height $z=h_{cars}=1.5$ m, of 7.48 $m^2 s^{-2}$.
369 We can combine this and the F_c value along with the area and volume of a lane of a roadway to estimate the energy density
370 (E_{VIT}) on dimensional grounds:

$$371 E_{VIT} = \left[\frac{(TKE)(air\ density)(lane\ volume)F_c}{(lane\ area)} \right] \quad (13)$$

372 Assuming each vehicle has a length of 4.5 m, width of 2.0 m, height of 1.5m, a lane length of 10 km, and an air density of
373 1.225 $kg\ m^{-3}$, one arrives at 84.8 $kg\ s^{-3}$, and values ranging from a North American grid maximum of 848 $kg\ s^{-3}$ to a background
374 value four orders of magnitude smaller (8.48×10^{-2} $kg\ s^{-3}$). These energy densities may be compared to the typical solar energy
375 density reaching the surface at mid-latitudes of 1300 $W\ m^{-2}$, or in SI units, 1300 $kg\ s^{-3}$, and the typical range of perturbations
376 in latent and sensible heat fluxes associated with the use of a more complex urban radiative transfer scheme (the Town Energy
377 Balance module; Mason, 2000) in our 2.5km grid cell size simulations (typical diurnal ranges in the perturbations associated
378 with/without use of TEB: latent: -200 to $+3$ $W\ m^{-2}$; sensible: -100 to $+100$ $W\ m^{-2}$ respectively). That is, under most daylight
379 conditions, the energy densities associated with VIT will be relatively small compared to the solar energy density at midday,



380 with a typical urban value of 6.5%, and range from 65% in the cell with the highest VKT values down to 0.0065% in
381 background conditions where the vehicle numbers are relatively small. Urban traffic however may contribute similar energy
382 levels as the changes in net latent and sensible heat fluxes associated with the use of an urban canopy radiative transfer model.
383 We also note that at night, during the low sun angle conditions of early dawn late evening, and during the lower sun angles of
384 winter, the relative importance of VIT to solar radiative input will be larger. Consequently, the impact of VIT will be higher
385 at night and in the early morning rush hour, and at other times when the sun is down or sun angles are low, as is demonstrated
386 below.

387 **2.7 GEM-MACH simulations**

388 A research version of the Global Environmental Multiscale – Modelling Air-quality and CHemistry (GEM-MACH) numerical
389 air quality model, based on version 2.0.3 of the GEM-MACH platform, was used for the simulations carried out here (Makar
390 *et al.*, 2017; Moran *et al.*, 2010; Moran *et al.*, 2018; Chen *et al.*, 2020). GEM-MACH is a comprehensive 3D deterministic
391 predictive numerical transport model, with process modules for gas and aqueous phase chemistry, inorganic particle
392 thermodynamics, secondary organic aerosol formation, vertical diffusion (in which area sources such as vehicle emissions are
393 treated as lower boundary conditions on the vertical diffusion equation), advective transport, and particle microphysics and
394 deposition. The model makes use of a sectional approach for the aerosol size distribution, here employing 12 aerosol bins.
395 The version used here also follows the “fully coupled” paradigm – the aerosols formed in the model’s chemical modules in
396 turn may modify the model’s meteorology via the direct and indirect effects (Makar *et al.*, 2015a,b; Makar *et al.*, 2017). The
397 meteorological model forming the basis of the simulations carried out here is version 4.9.8 of the Global Environmental
398 Multiscale weather forecast model (Cote *et al.*, 1998a,b; Caron *et al.*, 2015; Milbrandt *et al.*, 2016). Emissions for the
399 simulations conducted here were created from the most recent available inventories at the time the simulations were carried
400 out – the 2015 Canadian area and point source emissions inventory, 2013 Canadian transportation (onroad and offroad)
401 emissions inventory, and 2011-based projected 2017 US emissions inventory. As noted above, the model simulations were
402 carried out on two separate model domains shown in Figure 5; a 10 km horizontal grid cell size North American domain, and
403 a 2.5km horizontal grid cell size PanAm Games domain. For the 10km domain, simulations were for the month of July, 2016,
404 while for the higher resolution model, month-long summer (July 2015) and winter (January 2016) simulations were carried
405 out, with and without the VIT parameterization. These periods were based on the availability of emissions data, previous
406 model simulations for the same time periods appearing in the literature (Makar *et al.*, 2017; Stroud *et al.*, 2020), and the timing
407 of a prior field study (Stroud *et al.*, 2020).

408 **2.8 VIT as a Sub-grid-scale Phenomena**

409 It should be noted that the VIT enhancements to turbulent exchange coefficients are used to determine the vertical distribution
410 of freshly emitted pollutants at each model time step – they are not applied for all species within a model grid cell. Similar
411 sub-grid-scale approaches are used for the vertical redistribution of mass from large stack sources of pollutants, where



412 buoyancy calculations are applied to determine the rise and vertical distribution of pollutants from large industrial sources.
413 Both stacks and roadways are treated as sub-grid-scale sources of pollutants which are influenced by very local sources of
414 energy (stacks: high emission temperatures and exit velocities; roadways: vehicle induced turbulence) resulting in an enhanced
415 vertical redistribution of newly emitted chemical species. In both cases, the vertical transport results from an interplay between
416 the energy associated with the emission process (stacks: high temperature emissions with the ambient vertical temperature
417 profile; VIT: kinetic energy imparted to the atmosphere in which emissions have been injected with the ambient turbulent
418 kinetic energy). This interaction precludes a treatment solely from the standpoint of model input emissions, since the extent
419 of the mixing will depend on the local atmospheric conditions as well as the energy added due to the manner in which the
420 emissions occur. Both processes could be addressed by large eddy simulation modelling on a very local scale, but
421 parameterizations are required in both cases for regional scale simulations. In both cases, the parameterized vertical
422 redistribution of pollutants is applied to freshly emitted species – the horizontal spatial extent of the emitting region is
423 sufficiently small that although present, the enhanced mixing will have a minor effect on the redistribution of pre-existing
424 chemicals and on other atmospheric constituents affected by vertical transport. VIT in the context of regional chemical
425 transport models is thus best treated as a subgridscale phenomena applied to fresh emissions, in direct analogy to the approach
426 taken for large stack emissions.
427

428 **3 Results**

429 **3.1 VIT Height Dependence as a Gaussian Distribution**

430 Under Methods, we describe the potential for the use of a Gaussian distribution to describe the fall-off in TKE with height
431 above vehicles. Using the equations presented there, we have analyzed VIT studies appearing in the literature, determining
432 the decrease in TKE as a function of height from published figures, then fitting these data to a Gaussian distribution to the
433 height above ground. The result of this analysis for several data sets is shown in Table 1, generated by extracting vehicle
434 centerline TKE values from contour plots of published data, and is subdivided into isolated vehicle and vehicle ensemble
435 studies and cases.
436 The inferred mixing length shows a marked variation between that of isolated vehicles or the lead vehicle in an ensemble, and
437 that of other vehicles appearing further back in the ensemble. Both directly observed and CFD modelled values of the inferred
438 mixing length for *isolated* vehicles or the *lead* vehicles of an ensemble vary from 2.5 to 5.13 m. For subsequent vehicles in
439 an ensemble, the mixing lengths increase to range from 4.6 to 41 m. The difference in mixing length between the lead vehicle
440 in an ensemble, and subsequent identical vehicles appearing later in the ensemble also increases. For example note that diesel
441 truck mixing lengths inferred from the CFD modelling examining different vehicle configurations (Kim *et al.*, 2016) increase
442 from 5.13 to 14.64 m, and the mixing lengths for automobiles increase from 2.50 m (isolated automobile), to 4.6 m (automobile
443 two vehicles back from a lead diesel truck), to 9.41 m (automobile immediately behind a leading diesel truck). The mixing



444 length associated with VIT may also be significantly influenced by the ambient wind and local built environment – the mixing
445 length associated with the component of TKE due to VIT within street canyons (Woodward *et al.*, 2019; Zhang *et al.*, 2017)
446 ranges from 2/3 to greater than the street canyon height, with maximum mixing lengths of 41m. It is important to note that
447 these mixing lengths are driven by the vehicle passage within the canyon; they result from the additional TKE added
448 with/without vehicles in the CFD simulations. The above data show that a Gaussian distribution provides a reasonable
449 description of the decrease of TKE from vehicles with height, and, under realistic traffic conditions, the mixing lengths increase
450 in size, and are considerably larger than those of isolated vehicles, and are comparable to or greater than the near-surface
451 vertical discretization of air quality models.

452 The length scales associated with VIT range from 2.50 m in the case of isolated vehicles (Kim *et al.*, 2016), through ~10 m
453 for vehicles moving in ensembles (Woodward *et al.*, 2019; Zhang *et al.*, 2017) up to 41 m, with the larger values being typical
454 for urban street canyons. The latter describe the specific regions VIT is expected to have the greatest impact, given the high
455 vehicle density within the urban core. However, our parameterization makes use of length scales derived from observations
456 on open (non-street canyon) freeways (Gordon *et al.* 2012; Miller *et al.*, 2018), and thus may underestimate the length scales
457 in the urban core. The impact of multiple vehicles travelling in an ensemble on open roadways was specifically depicted in
458 the open roadway simulations of Kim *et al.* (2016) reproduced in Methods (Figure 1), where the vertical extent of turbulent
459 mixing was shown to grow with increasing number of vehicles travelling in an ensemble. Furthermore, as was discussed and
460 demonstrated in Methods using the diffusivity equation, the length scale of the turbulence need not be greater than the model
461 lowest layer resolution in order to capture the impacts of VIT on mixing, being due in part to the gradient in turbulence with
462 height.

463 3.2 Model Domains and Evaluation Data

464 Our 3D air-quality model (GEM-MACH) and our VIT parameterization, including its diurnal variation, are described under
465 Methods. Two air-quality model grid cell size and domain configurations were used for our simulations – the first employs a
466 10km grid cell size with a North American domain, and is used for the current operational GEM-MACH air-quality forecast
467 (Moran *et al.*, 2010; Moran *et al.*, 2018; Figure 5(a)). The second was a 2.5km grid-cell resolution domain focused on the
468 region between southern Ontario, Quebec and northeastern USA (Joe *et al.*, 2018; Ren *et al.*, 2020; Stroud *et al.*, 2020; Figure
469 5(b)).

470 The impact of VIT was determined through paired model simulations, with and without the VIT parameterization, evaluated
471 against surface monitoring network data. The latter include hourly model output for ozone (O₃), nitrogen dioxide (NO₂), and
472 particulate matter with diameters less than 2.5 μm (PM_{2.5}), across North America and in our high resolution eastern North
473 America domain, evaluated at observation station locations with data from the AirNow network (AirNow, 2020). Observation
474 station locations used in simulation evaluation for these species are shown in Figure 6, for the two model configurations. The
475 juxtaposition of observation stations with urban populations (where the highest vehicle density may be found) may be seen by
476 comparing Figure 6 with Figure S2.



477 3.3 Continental 10km Grid Cell Size Domain Evaluation

478 Simulations were carried out for the month of July, 2016 for the 10km grid cell size North American domain. Model
479 performance metrics used to here (see Methods) are described in Table S1, and provided for the 10 km resolution “VIT” and
480 “No VIT” simulations relative to the hourly observation data for PM_{2.5}, NO₂, and O₃ in Table 2. These three chemicals were
481 chosen due to their well-known link to human health impacts of air pollution (Steib *et al.*, 2008; Abelson *et al.*, 2011).

482 The addition of VIT improved the scores for most performance metrics (bold-face print in Table 2). For NO₂, the addition of
483 VIT improved all scores with the exception of the correlation coefficient, which was degraded in the third digit. All PM_{2.5}
484 scores improved, with the exception of the mean bias, which became more negative by 0.5 µg m⁻³ across North America. All
485 ozone scores improved, the exceptions being the correlation coefficient (which was the same for both simulations, or improved
486 in the 3rd digit depending on the domain or country), and the ozone mean bias for the USA (which increased by +0.18 ppbv).
487 Some of the improvements were substantial, when considered in a relative sense: this was most noticeable for the NO₂ scores,
488 with the North American Mean Bias for NO₂ improving by a factor of 8.4, the mean gross error and index of agreement by
489 19%, the root mean square error by 25%, and the FAC2 score by 6%. Relative improvements for PM_{2.5} across North America
490 were more modest (ranging from 0.3% for FAC2 to 14% for the correlation coefficient. The corresponding relative changes
491 for O₃ ranged from a 22% reduction in the mean bias magnitude to a fraction of a percent improvement for FAC2, mean gross
492 error, root mean square error, and index of agreement. Overall, the model performance for the Continental 10km domain July
493 2016 simulations improved across different metrics, indicating that the increased vertical turbulent mixing resulting from the
494 incorporation of VIT results in a more accurate representation of atmospheric mixing and chemistry.

495 The timing and spatial distribution of the differences in the 29 day mean values of NO₂, PM_{2.5} and O₃ at 10 and 22 UT (6 AM
496 and 6 PM EDT) are shown in Figure 7. NO₂ and PM_{2.5} have decreased in the urban areas and along the major road networks
497 in the early morning (Figure 7(a,b)), while the ozone (Figure 7(c)) increases in the urban areas and along the roadways, with
498 a minor increase in the surrounding countryside. The VIT effect occurs at night and in the early morning: the average
499 differences are minimal by 6 PM EDT (Figure 7 (b,d,f)). This diurnal cycle of the average impact of VIT is expected: at night
500 and during the early morning the radiative-transfer driven atmosphere is relatively stable, natural background turbulence is
501 low in magnitude, and the relative contribution of VIT is therefore large. The reverse is true during the later morning to late
502 afternoon, as the solar radiative balance causes near-surface turbulence to rise several orders of magnitude relative to nighttime
503 values, and the relative contribution of VIT at those times becomes minimal. The strongest contribution of VIT thus occurs
504 under more stable atmospheric conditions: at night and in the early morning.

505 The all-domain model performance metrics of Table 2 were also calculated for each measurement station, and the appropriate
506 differences in the metrics or their absolute values were used to determine location-specific impacts of the VIT parameterization
507 for NO₂, PM_{2.5} and O₃ (Figures 8, S3 and S4). Differences in the values of the metrics between the two simulations are
508 shown, with the sign of the differences arranged so that red/blue colours indicate better performance for the VIT/No VIT
509 simulations respectively, red indicating better scores for the VIT simulation. The colour scales in these Figures are arranged



510 to include 3 orders of magnitude between lowest and highest difference scores and zero, and to encompass the maximum value
511 of the differences observed at across all stations. The values vary between metrics and the chemical species, with the largest
512 changes occurring for NO₂, followed by PM_{2.5} and the smallest changes for O₃, relative to typical concentrations of these
513 species, and in accord with Table 2. NO₂ performance improvements with the VIT simulation (red colours) occur across most
514 stations for the FAC2, MGE, RMSE, COA and IOA scores (Figure 8(a,c,e,f,g)), while r and |MB| scores are more variable,
515 with some stations having better performance for the No VIT simulation. PM_{2.5} performance improvements are more mixed,
516 with large improvements for correlation coefficient (Figure S3(d)) and IOA (Figure S3(g), a mild but overall positive effect of
517 VIT for MGE, RMSE and COE (Figure S3(c,e,f)), and more stations showing a degradation of performance for FAC2 and
518 |MB|, echoing the net effect for these last two metrics seen in Table 2. O₃ performance shows a strong regional variation
519 (Figure S4): most scores improve with the use of the VIT parameterization in the western and north-eastern parts of the
520 continent, and degrade in the south-eastern USA. The degradation in the south-eastern (e.g. increases in O₃ concentrations in
521 a region which already experiences a positive O₃ bias) are associated with the transport of urban O₃ precursors into forested
522 areas in the region, with additional O₃ production occurring there. These effects may be removed through the introduction of
523 an additional parameterization for the reduced turbulence and shading within forested canopies (Makar *et al.*, 2017; Figure
524 S5), with the combined parameterizations resulting in improvements in both NO₂ and O₃ performance. While the use of VIT
525 degrades O₃ performance in this region, this degradation is thus very small relative to the large improvements noted with the
526 canopy effect (see Makar *et al.*, 2017; Fig S5 and its associated discussion in the S.I.). Another significant feature is the
527 improvement (red colours) in most O₃ station scores in urban regions (Figure S4). These improved scores largely result from
528 increases in ozone in the early morning hours (Figure 7(e)), where VIT has resulted in increased vertical mixing, reducing
529 surface level NO_x and hence NO_x titration of ozone, and also by mixing higher ozone levels aloft down into the lowest model
530 layer.

531 Overall, the impact of the VIT parameterization was to improve North American simulation accuracy, across multiple
532 statistical metrics, with the most significant improvements in the model performance for simulated NO₂. Spatially, model
533 performance was generally greatest in urban regions and western and northeastern North America, though this depends on the
534 chemical species and the performance metric chosen.

535 3.7 Eastern North America 2.5km Grid Cell Size Domain Evaluation

536 With the use of a smaller grid cell size (i.e. “higher resolution”), meteorological models and on-line air-quality models such
537 as GEM-MACH have the option of employing theoretical approaches which better simulate the more complex radiative
538 transfer and physical environment-induced turbulence of urban areas. Urban heat islands are known to have a significant effect
539 on turbulence, for example (Mason, 2000; Makar *et al.*, 2006). In these simulations, we make use of the Town Energy Balance
540 (TEB; Mason, 2000; Leroyer *et al.*, 2014; Lemonsu *et al.*, 2005), a single-layer urban canopy module which solves the
541 equations for urban atmosphere’s surface and energy budgets for a variety of urban elements (roads, walls, roofs), then
542 aggregates the results for the net urban canopy. Such parameterizations are inappropriate for use in larger grid cell size models



543 due to the latter's inability to resolve individual surface types and spatial gradients at the city scale. An important consideration
544 in determining the relative importance of vehicle-induced turbulence is whether improvements in performance still occur,
545 when these other sources of turbulent kinetic energy are included explicitly. We address this issue in our 2.5km grid cell size
546 modelling by employing the TEB parameterization, for both VIT and No VIT simulations, evaluating both simulations against
547 surface monitoring network observations as before. Both summer and winter simulations were carried out on the blue domain
548 of Figure 5(b), and the same performance metrics were calculated as for the larger North American simulations (Table 3).
549 A similar pattern of performance improvement can be seen between 10km and 2.5km grid cell sized, comparing Tables 2 and
550 3, with improvements due to the use of VIT predominating in both summer and winter: despite the addition of a more explicit
551 urban radiative balance approach, better scores were achieved with the addition of the VIT parameterization. Note that
552 comparisons between the 2.5km and 10km simulations for similar emissions inputs appear elsewhere in the literature (Stroud
553 *et al.*, 2020). The number of improved scores increases from summer to winter. Stable atmospheric conditions and low
554 meteorological turbulence levels are more common in winter than summer, during both day and night, and the impact of the
555 additional source of turbulence is thus proportionally stronger in the winter season. The VIT effects at the urban scale are the
556 strongest for NO₂ and PM_{2.5}, and less noticeable for simulated O₃, similar to the North American domain simulation. The
557 largest improvements for the three species and across seasons occur for winter PM_{2.5}, with the improved performance taking
558 place in the first or second digit of the given metric. Metric differences for NO₂ aside from mean bias occur in the second to
559 third digit in the winter, with summer differences occurring in the first to 2nd digit. Changes to O₃ are relatively minor, with
560 some improvements and degradation in performance in the 3rd digits across the different metrics.

561 UT-hour average differences between the two 2.5km grid-cell size simulations, for the three species evaluated for the summer
562 and winter simulations, appear in Figures (S6, S7), and Figures (9, 10) respectively. The summer differences in surface
563 concentration (Figure S6) are the largest at 6 AM local time (10UT; first column of panels), and have largely decreased to near
564 zero by 6 PM (22 UT; last column). Corresponding concentration vertical distribution differences along a cross-section linking
565 the major cities show the early morning depletion (increase) of NO₂, PM_{2.5} (O₃) are coupled to increases (decreases) aloft
566 (Figure S7, first column of panels). NO₂ and PM_{2.5} reductions extend to altitudes of up to 2km with the increase in radiative-
567 driven turbulence during the day, while the change in NO_x/VOC regime aloft leads to increases in lower Troposphere O₃
568 (Figure S7, second column). Daytime mixing increases lead to a reduction in the effect by nightfall (Figure S7, third column).
569 VIT-enhanced transport of NO₂ from urban to rural areas can also be seen (Figure S6, center column/first column; note
570 increases in NO₂ on the periphery of the urban areas, pink to red colours). This additional NO_x added to NO_x-limited regions
571 leads to low-level (mostly sub-ppbv) increases in daytime O₃ at 10AM which persist through to 6PM. Over the Great Lakes,
572 the change in vertical transport on land, coupled with daytime lake breeze circulation (Makar *et al.*, 2010; Joe *et al.*, 2018;
573 Stroud *et al.*, 2020) results in a decrease in daytime NO₂ and PM_{2.5} over the Lakes and corresponding late-afternoon O₃
574 increases (Figure S6, blue colours in centre column of panels over the lakes for NO₂ and PM_{2.5}, red colours in the final panel
575 of the sequence for O₃). The changes in the near-roadway environment thus have larger regional effects, changing the pathway
576 and reaction chemistry of transported chemicals on a regional scale.



577 The stronger impact of VIT under winter conditions is illustrated in Figures 9 and 10; NO₂ decreases (Figures 9,10 (a,b,c))
578 persist throughout the day, though to a lower degree by 6 PM (contrast Figures S6,S7 (a,b,c) to Figures 9,10(a,b,c)). The
579 vertical influence of VIT reaches an altitude of approximately 2 km in the winter (1 km in the summer); contrast Figure S7
580 and Figure 10. The absence of winter biogenic hydrocarbon production during the day has likely limited the daytime increase
581 in O₃ to the cities (compare Figure S6(h) with Figure 9(h)). The large effect of VIT along major roadways can be seen in
582 both Figures S6 and 9, particularly in the 6AM column of panels (a,d,g) in both figures, with the greatest reductions aside from
583 urban regions occurring along major roadways (e.g. Chicago to Detroit area).

584 Differences in station-specific performance scores for the two simulations for the 2.5km grid-cell size domain, constructed as
585 for the 10km domain, are shown in Figures S8, S9, and S10 (summer) and Figures S11, S12 and S13 (winter) for NO₂, PM_{2.5}
586 and O₃, respectively.

587 The summer scores (Figs. S8, S9, S11) show the most significant improvements in the urban areas across all performance
588 metrics, with the largest relative magnitude differences for NO₂ and PM_{2.5}, and lower magnitude changes for O₃. As for the
589 North American simulations, O₃ performance improvements occur in the cities, due to increased vertical mixing, and, O₃ scores
590 in rural regions have degraded, but may be improved with the use of a forest canopy parameterization, as discussed further in
591 the SI (Figure S10, S13, S5). The overall impact of the incorporation of the VIT parameterization is clearly a positive one,
592 particularly in urban areas: VIT has been shown to have a significant impact on summertime urban and suburban scale
593 photochemistry.

594 The metrics of the winter 2.5km station-specific evaluation for NO₂ (Figure S11) show both local improvements and
595 degradation in performance, depending on location. Wintertime PM_{2.5} performance improves across most metrics and most
596 locations (Figure S12). Wintertime ozone performance is variable, though improvements can be seen for most metrics within
597 the largest urban areas (Figure S13).

598 **4 Discussion and Conclusions**

599 Our work implies that the turbulence associated with vehicle motion is capable of having a significant effect on the
600 concentrations of key pollutants in the lower atmosphere. Incorporating that effect into both continental-scale and higher
601 resolution regional/urban scale air implementations of a pollution model resulted in an overall improvement in model
602 performance, across several different performance metrics. The improvement at higher resolution (when the TEB urban
603 parameterization was included in the model setup) implies that the mixing associated with urban radiative transfer and
604 roughness is not sufficient to account for the observed pollutant concentrations; the effect of VIT is robust despite differences
605 in radiative transfer schemes and across different horizontal resolutions.

606 However, we also acknowledge several limitations of our VIT formulation and have recommendations for future work which
607 would allow it to be improved and the uncertainties in our analysis reduced.



608 First, we have assumed that single-vehicle induced turbulence accounts for all of the turbulent kinetic energy contributed by
609 vehicles (Gordon *et al.*, 2012; Miller *et al.*, 2018). The passage of multiple vehicles also induces a “wake flow” in their
610 direction of motion. While this effect has been recognized in very high resolution roadway-scale models (Eskridge and
611 Catalano, 1987; Eskridge *et al.*, 1991), the breakdown of opposing wake flows into turbulence (arising from two-way traffic
612 and/or multiple lanes of traffic travelling at different speeds) has not been examined, to the best of our knowledge. However,
613 these wake flows are of sufficiently high energy that their residual power is being harnessed via vertical-turbine wind power
614 generation systems in both Turkey (Devecitech, 2020) and Scotland (Shell, 2020). The single-vehicle additive
615 parameterization we have created here may thus underestimate the net turbulent effect of vehicle passage. At the same time,
616 our assumption that individual VIT within a grid cell is simply additive may also be incorrect, resulting in overestimates of
617 that portion of the net VIT. With the advent of Doppler LIDAR systems with sufficient time resolution to capture turbulence,
618 we advocate for and are currently embarking on new observation studies employing these systems in scan mode across
619 highways, to fully characterize all vehicle-induced contributions to turbulence as a function of the number and type of vehicles
620 crossing below a LIDAR scan path perpendicular to the highway.

621 Second, our assumption that each vehicle’s pathway crosses the grid cell is a considerable source of uncertainty. There we are
622 limited by the lack of availability of simultaneous vehicle speed and number data. However, recent developments in satellite-
623 based radar technology have been shown to provide accurate estimates of the speed of individual vehicles along major
624 highways (Meyer *et al.*, 2006; Bethke *et al.*, 2006), and binning and collection of these data may improve the linkage between
625 the more commonly available vehicle-km-travelled data and VIT beyond that used here. Other sources of gridded vehicle
626 and/or road density data (World Bank, 2018) should also be explored.

627 Third, one consideration for our parameterization is the issue of “traffic jams”; a large number of vehicles being present on the
628 road without much motion in such conditions. However, we note that in this case, the number of vehicles crossing a point in
629 space will drop – that is, if the underlying traffic data (vehicle-km-travelled) is of sufficient quality that traffic jams are
630 included, the existing parameterization should adequately handle these effects. Both our second and this third consideration
631 argue for the creation of more accurate vehicle travel data for use in air-quality models.

632 Last, we note that the ambient concentrations of pollutants such as NO₂, O₃ and PM_{2.5} are influenced by a host of factors
633 included in other parameterizations of air-quality models, and in the quality of the available emissions data. However, we
634 have shown here that improvements in the forecast quality of three different species with human-health impacts may be
635 achieved through the same process improvement. An examination of all of the other possible sources of error in air-quality
636 models is beyond the scope of this work. We also emphasize that the work does not identify a deficiency in existing
637 meteorological boundary layer turbulence models. Rather, that the at-source vertical transport of fresh pollutants from the
638 mobile sector needs to take into account local sources of energy for transport at the point of emission (whether in large stacks
639 (Gordon *et al.*, 2018; Akingunola *et al.*, 2018) or over roadways (as examined here)).

640 Despite the uncertainties identified above, our analysis has shown:



641 (1) The drop-off of VIT with height above moving vehicles is well-represented by a Gaussian distribution, from multiple
642 measurement and computational fluid dynamics modelling studies.

643 (2) The mixing lengths inferred from these studies ranges from 2.50 m (for individual isolated cars) through ~10 m
644 (vehicle ensembles) to 41 m (vehicle ensembles in a street canyon environment). We also note that the gradient in the net
645 thermal turbulent transfer coefficients drives concentration changes due to VIT. The expectation that VIT is capable of vertical
646 transport out of the lowest layers of a regional model is therefore a reasonable one.

647 (3) The magnitude of the localized energy input from VIT, while smaller than the input of solar energy during daylight
648 hours, is equivalent in magnitude to the energy perturbations resulting from the use of a state-of-the-art urban radiative balance
649 model (TEB; see Methods). That is, locally, VIT has sufficient energy to be equivalent to the impact of an improved urban
650 radiative transfer scheme – underlining its importance for vertical transport of pollutants.

651 (4) The impact of VIT depends on both local traffic conditions and the background meteorological conditions, with the
652 maximum effect occurring when turbulence in the ambient atmosphere is relatively weak (night through early morning), and
653 traffic levels are relatively high (morning rush hour).

654 (5) The use of the VIT parameterization has been demonstrated to result in decreases in air-quality model error, across
655 three different key pollutants, with the most striking results for mean biases, without resorting to the use of imposed minima
656 in the thermal turbulent exchange coefficients frequently used in air-quality models.

657 (6) VIT has a significant impact on the rapid vertical distribution of freshly emitted pollutants on the very localized scale
658 of roadways where the enhanced mixing occurs, in analogy to the rapid vertical transport used in parameterizations of plume
659 rise from large stacks. Its impact on mixing of pre-existing meteorological and chemical variables on the grid-cell scale is
660 expected to be small.

661 Based on these findings, we conclude that VIT has a significant impact on pollutant transport and dispersion out of the lowest
662 layer of the atmosphere, and recommend its inclusion in air-quality models. Further improvements to the parameterizations
663 found herein would result from additional observations of TKE using Doppler lidar techniques, of vehicle ensembles under
664 realistic driving conditions.

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666

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668 of VKT gridded fields.

669

670 **References**

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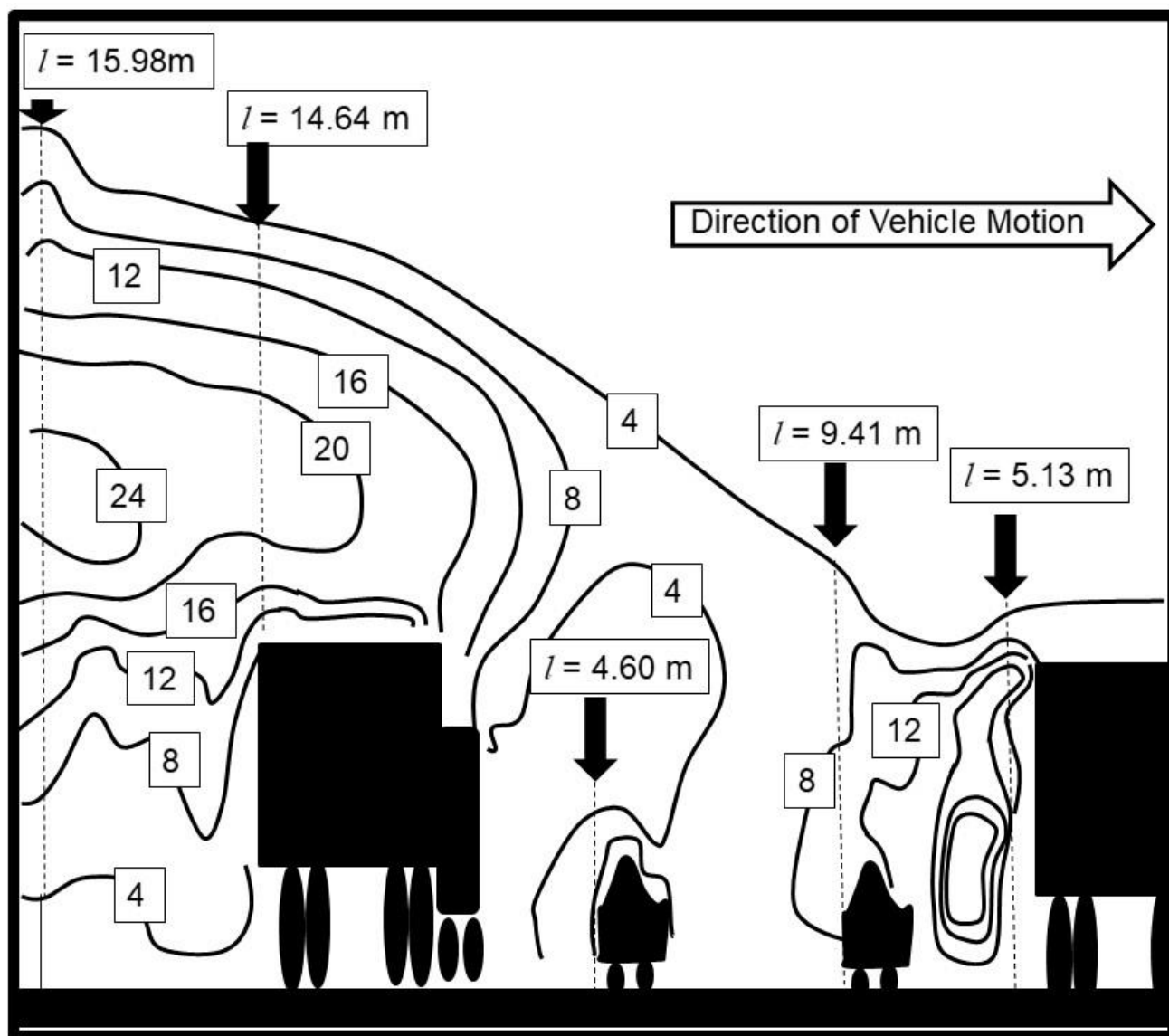
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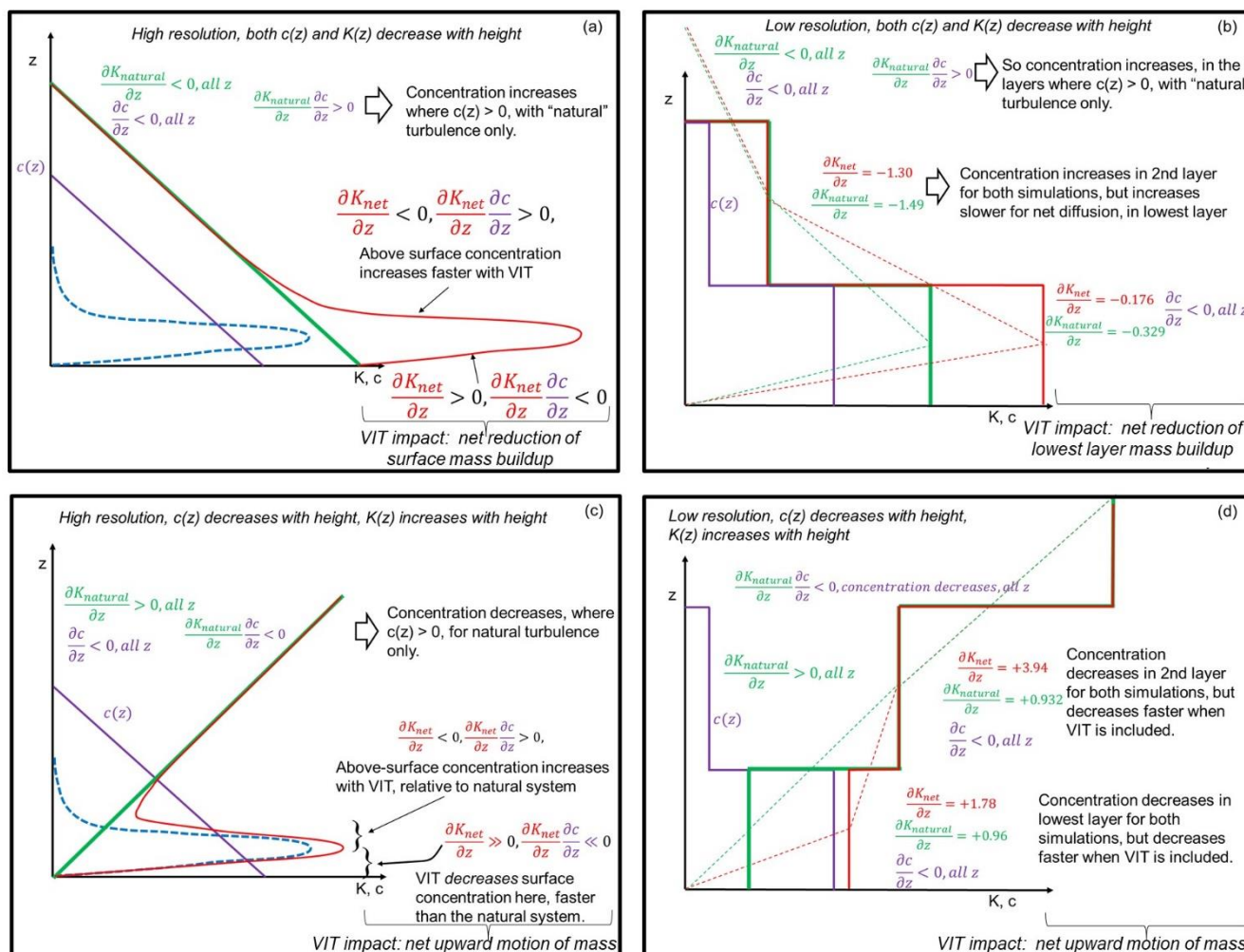
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830 **Figure 1.** Example of length scales associated with an ensemble of vehicles (after Kim *et al.*, 2016, Figure 14). TKE contours along dashed
831 lines were extracted and fit to equations (1,2) for Table 1. Note that the length scale of turbulence immediately behind the leading vehicle,
832 a large transport truck is only 3m, while the length scale immediately behind the trailing vehicle in the ensemble (an identical transport truck)
833 is 12.73m.

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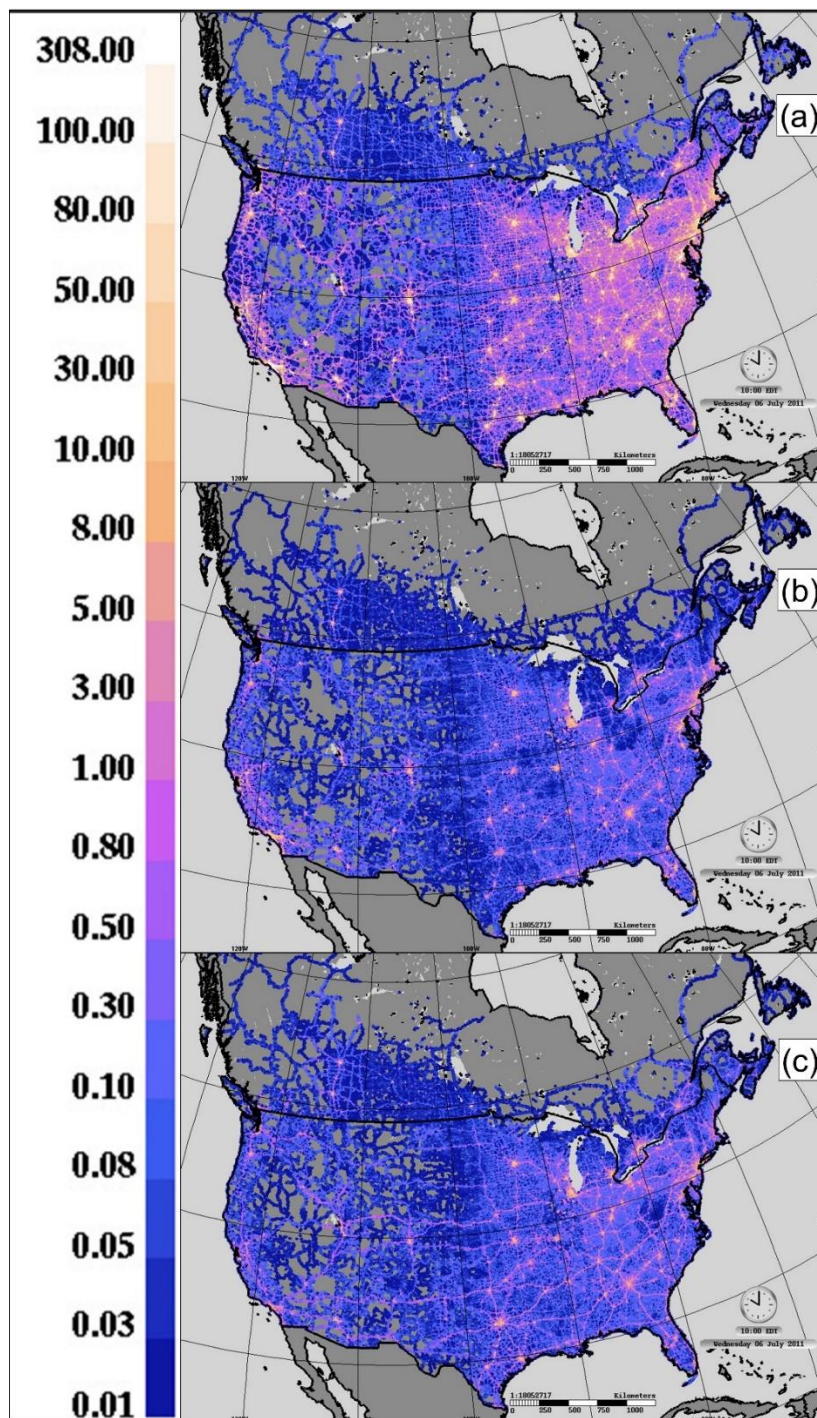
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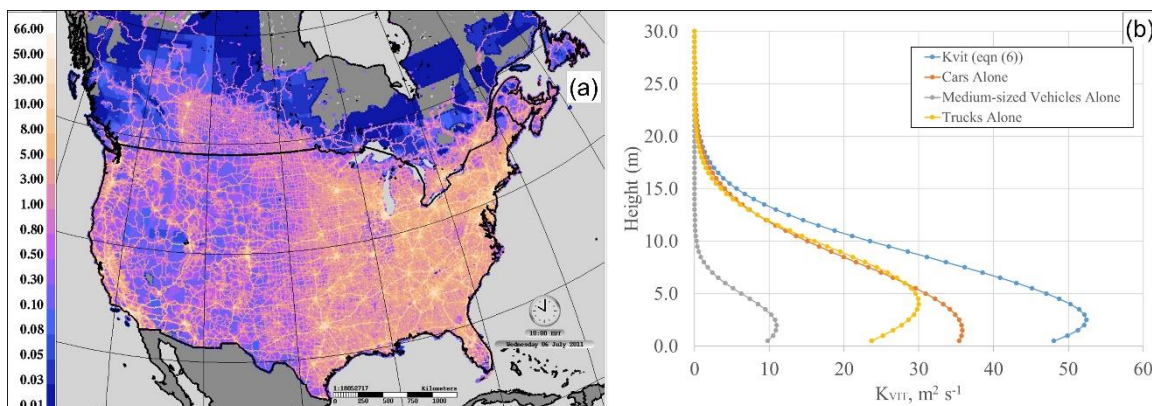
Figure 2. Illustration of the impact of VIT on the local vertical gradient of the thermal turbulent transfer coefficients, at low (a,c) and high (c,d) resolution. Purple, green, dashed blue, and red lines illustrate the height variation of concentration, meteorological or natural coefficient of thermal turbulent transfer, VIT coefficient of thermal turbulent transfer, and net coefficient of thermal turbulent transfer, respectively. (a,b) High and low resolution profiles and gradients, for the case where both concentration and meteorological thermal turbulent transfer coefficients decrease with height. (c,d) High and low resolution profiles and gradients, for the case where concentration and meteorological thermal turbulent transfer coefficients increases with height.



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843 **Figure 3.** Vehicle km travelled per 10 km grid cell (km s^{-1}) for (a) cars, (b) mid-size vehicles and (c) trucks, July, 2015.

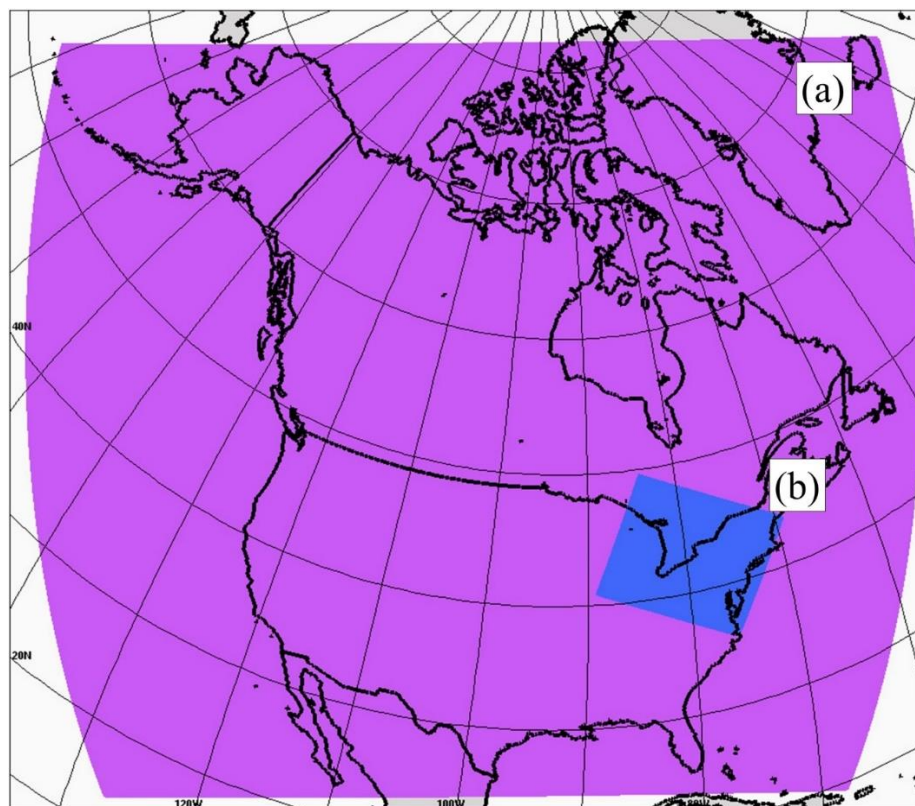
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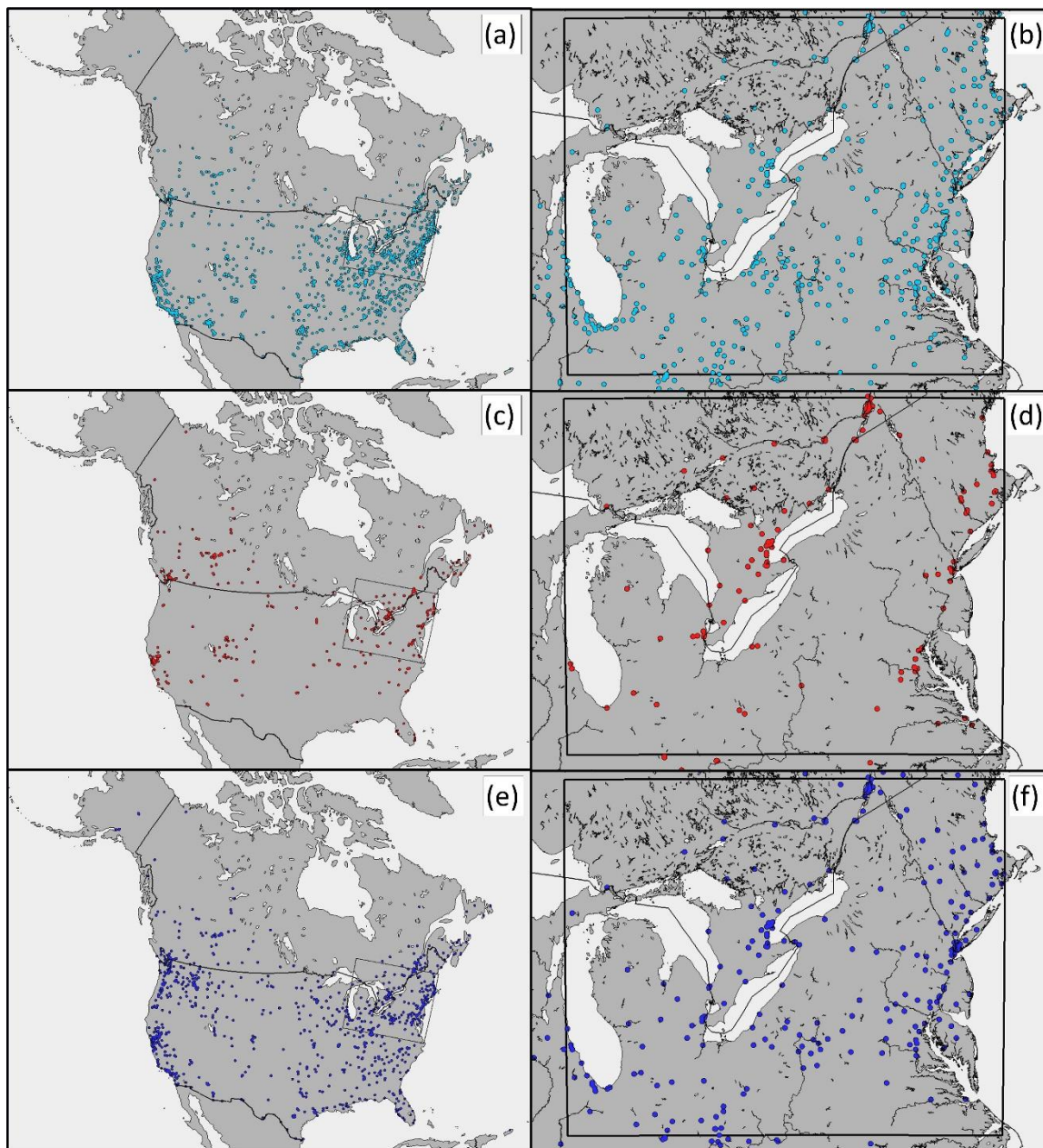
846 **Figure 4.** (a) Example estimated thermal turbulent transfer coefficients from VIT at 2 m elevation during a weekday at 10 am in July (m^2s^{-1}), using the VKT data of Figure 3. (b) Vertical profile of VIT thermal turbulent transfer coefficients at one meter resolution in central
847 Manhattan Island, and individual values for the TKE associated with cars, mid-sized vehicles and trucks considered separately, generated
848 using equation (8). Note that the profiles of (b) would be added to the ambient thermal diffusivity coefficients (see section 2.5, and
849 equation (12)).
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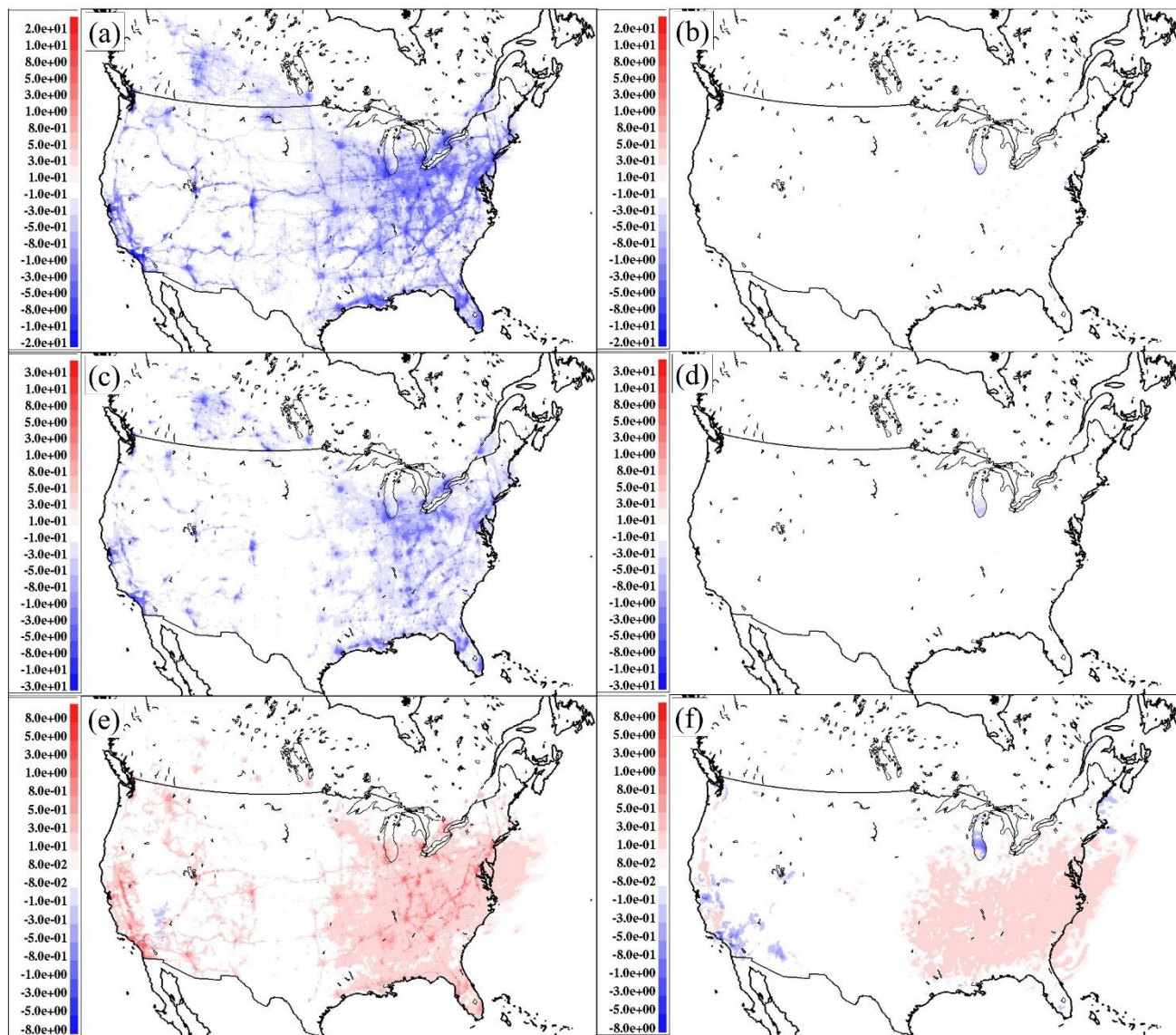
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853 **Figure 5.** GEM-MACH test domains: (a) 10km grid cell size North American domain. (b) 2.5km grid cell size Pan Am domain.
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856 **Figure 6.** AIRNOW hourly observation station locations for ozone (a,b), nitrogen dioxide (c,d), and particulate matter with diameters less
857 than $2.5 \mu\text{m}$ (e,f). (a,c,e): Stations used for the 10km grid cell size domain evaluation. (b,d,f): Stations used for the 2.5km grid cell size
858 domain evaluation (all stations located within central box).

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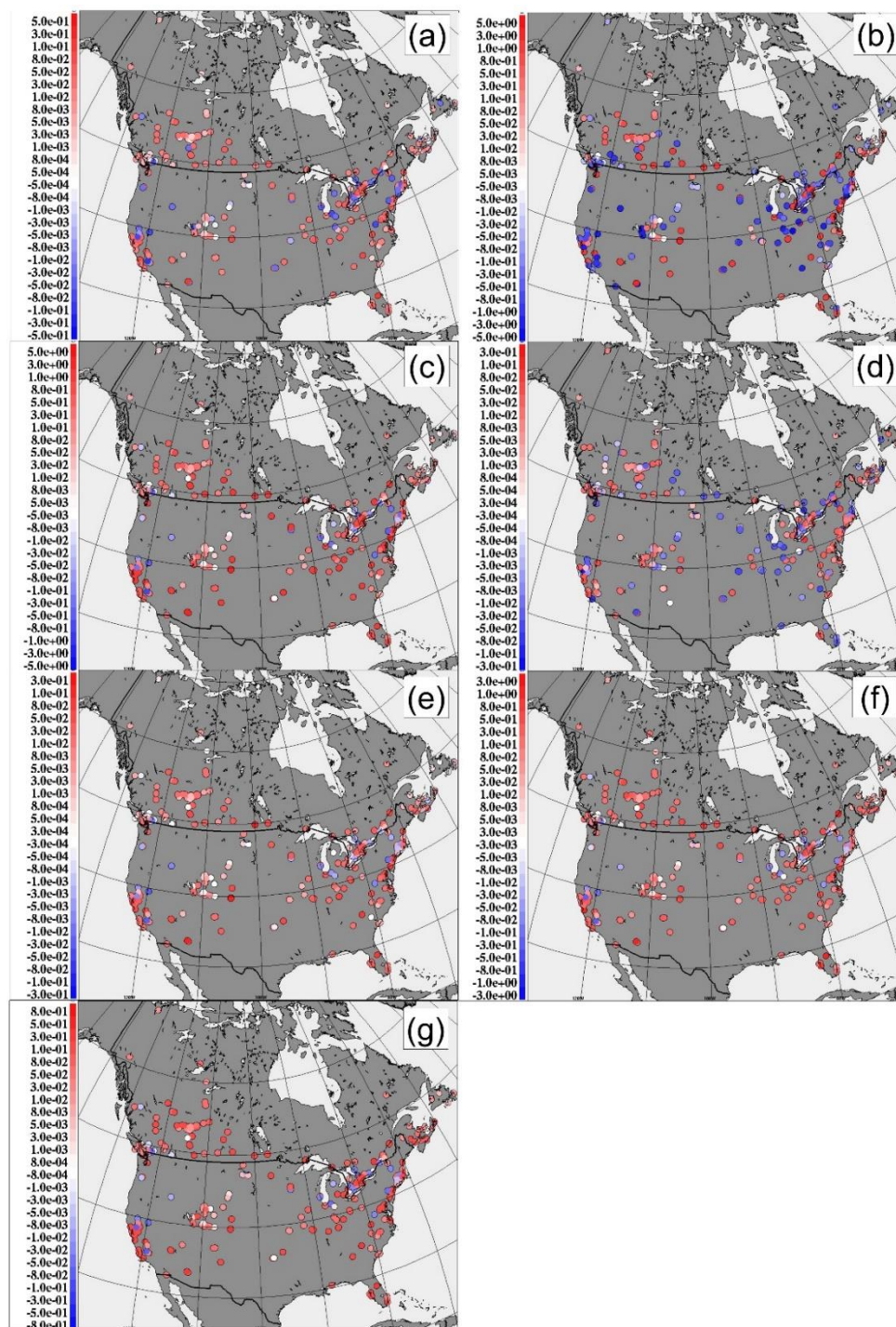


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863 **Figure 7.** Difference in 29 day average NO₂, PM_{2.5} and O₃, July 2016 Continental 10km domain simulations (VIT simulation – No VIT
864 simulation). Averages are paired at (a,c,e) 10UT, b,d,f) 22UT according to species; (a,b): ΔNO₂(ppbv); (c,d) ΔPM_{2.5}(μg m⁻³); (e,f)
865 ΔO₃(ppbv).

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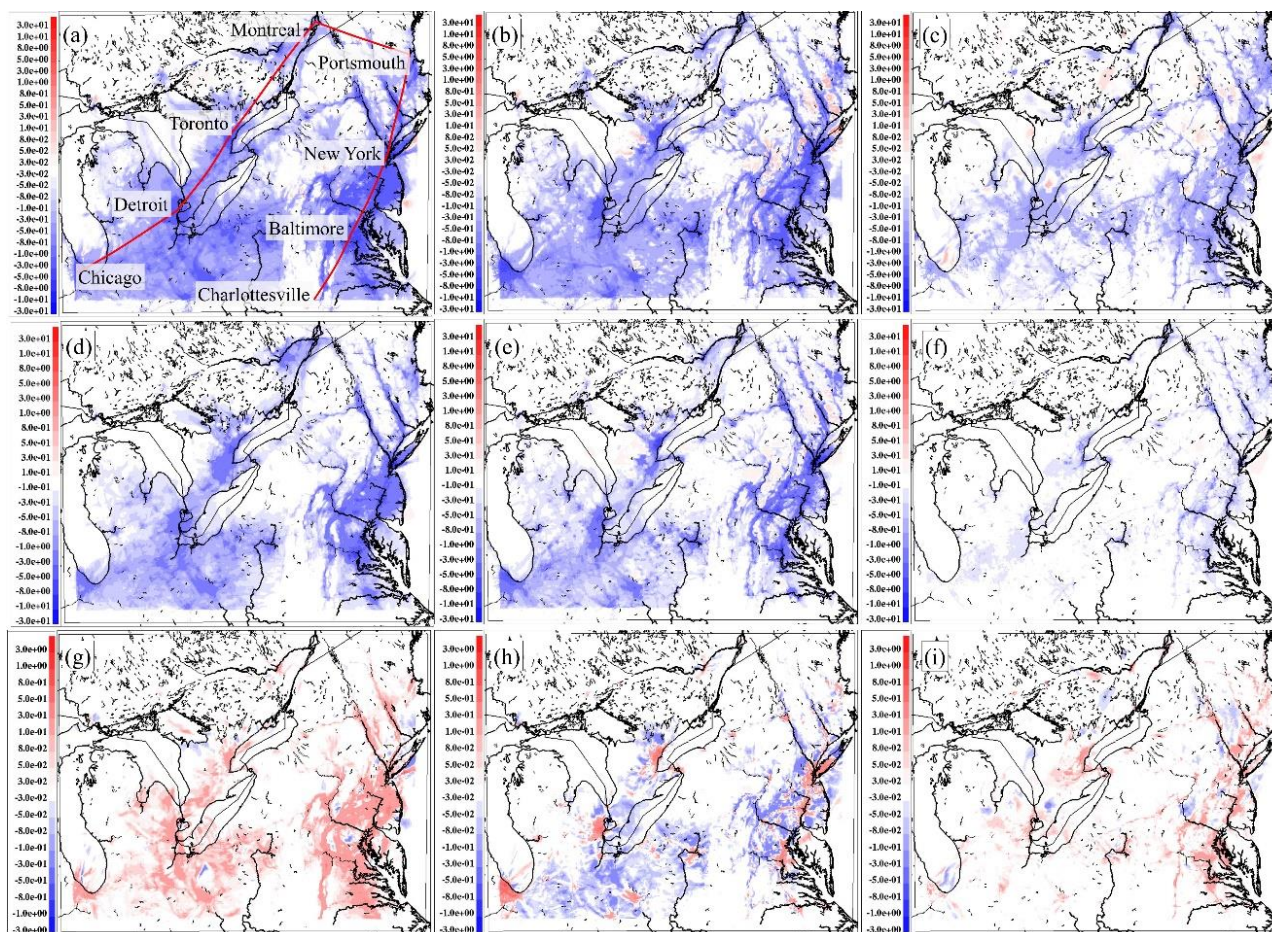
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869 **Figure 8.** Change in model NO₂ performance at 358 North American surface monitoring sites, July 2016 (ppbv). Red colours indicate
 870 stations where the addition of the VIT parameterization improved model performance, blue colours indicate stations where the addition of
 871 the VIT parameterization degraded model performance. (a) $\Delta FAC2_{No\ VIT-VIT}$; (b) $\Delta |MB|_{No\ VIT-VIT}$; (c) $\Delta MGE_{No\ VIT-VIT}$; (d) $\Delta r_{No\ VIT-VIT}$;
 872 (e) $\Delta RMSE_{No\ VIT-VIT}$; (f) $\Delta COE_{No\ VIT-VIT}$; (g) $\Delta IOA_{No\ VIT-VIT}$.



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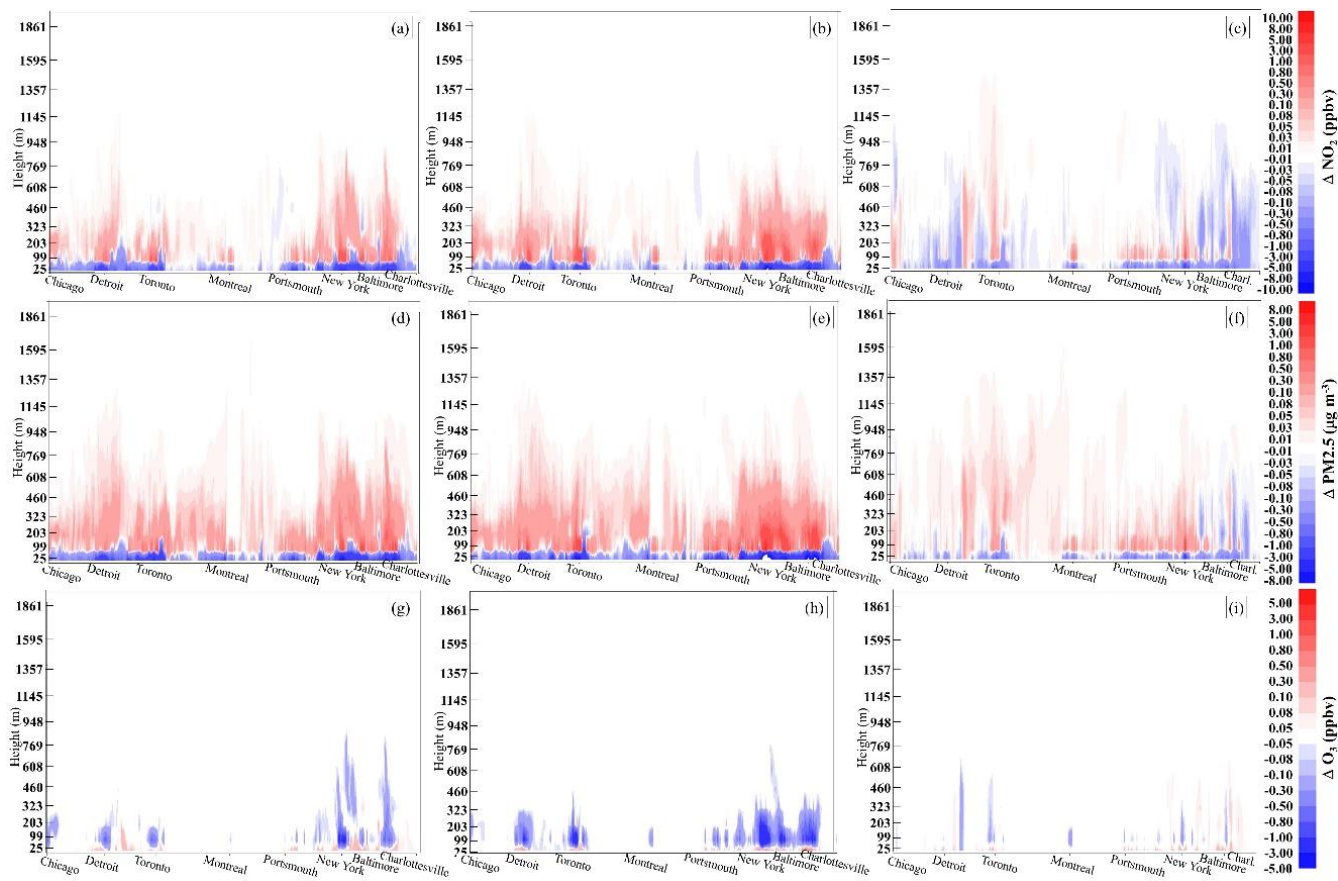
Figure 9. Difference in 30 day average surface NO_2 , $\text{PM}_{2.5}$ and O_3 , January 2016, PanAm 2.5km grid cell size domain simulation. Averages are paired at (10, 14, and 22UT) according to species; (a,b,c): ΔNO_2 (ppbv) (d,e,f) $\Delta\text{PM}_{2.5}$ ($\mu\text{g m}^{-3}$); (g,h,i) ΔO_3 (ppbv). Red line in panel (a) indicates position of vertical cross-section shown in Figure 10.

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883 **Figure 10.** Vertical cross-sections of concentration differences between major eastern North American cities, January 2016, panels arranged
884 as in Figure 9. Vertical coordinate: unitless hybrid, top-of-scale is approximately 2 km.

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Table 1. Gaussian distribution fits of VIT TKE drop-off with height, from observation and CFD studies.

Study, Case	Slope	Intercept	R ²	Mixing length $(z \text{ at } e^{\left(\frac{(z-h_q)^2}{2\sigma_q^2}\right)} = 0.01), \text{ m}$
<i>Isolated vehicles:</i>				
Rao <i>et al.</i> (2002), cube van, 50 mph, $h_q = 2\text{m}$	2.2452	1.8534	0.9856	3.53
Rao <i>et al.</i> (2002), cube van, 30 mph, $h_q = 2\text{m}$	1.0230	1.4969	0.9709	4.22
Kim <i>et al.</i> (2016), lead automobile, $h_q = 1.5\text{m}$	4.6431	3.9013	0.8845	2.50
Kim <i>et al.</i> (2016), lead diesel cargo truck, $h_q = 4\text{m}$	3.6143	4.2223	0.9355	5.13
<i>Vehicle Ensembles:</i>				
Kim <i>et al.</i> (2016), automobile immediately following lead diesel cargo truck, $h_q = 1.5\text{m}$	0.073529	4.1144	0.9801	9.41
Kim <i>et al.</i> (2016), 2 nd automobile, following lead diesel cargo truck, $h_q = 1.5\text{m}$	0.47337	3.9275	1.00 ^a	4.60
Kim <i>et al.</i> (2016) 2 nd diesel cargo truck, $h_q = 4\text{m}$	0.04070	4.7935	0.5424	14.64
Woodward <i>et al.</i> (2019) vehicle ensemble ^b , $h_q = 1.5\text{m}$, parallel to flow, right lane	0.01916	-1.2402	0.9135	17.01
Woodward <i>et al.</i> (2019) vehicle ensemble ^b , $h_q = 1.5\text{m}$, parallel to flow, left lane	0.01155	-1.4532	0.7543	21.46
Woodward <i>et al.</i> (2019) vehicle ensemble ^b , $h_q = 1.5\text{m}$, transverse to flow, right lane	0.012489	-1.4766	0.9667	20.70
Woodward <i>et al.</i> (2019) vehicle ensemble ^b , $h_q = 1.5\text{m}$, transverse to flow, left lane	0.0098094	-1.7815	0.9536	23.16
Zhang <i>et al.</i> (2017), VS1: $h_q = 1.6\text{m}$, vehicle speed = 9 km hr ⁻¹ , Wind 11 km hr ⁻¹	0.0029165	5.1706	0.6614	41.24
Zhang <i>et al.</i> (2017), VS2: $h_q = 1.6\text{m}$, speed = 36 km hr ⁻¹ , Wind 11 km hr ⁻¹	0.005158	5.0964	0.8306	31.38
Zhang <i>et al.</i> (2017), VS3: $h_q = 1.6\text{m}$, vehicle speed = 36 km hr ⁻¹ , Wind 36 km hr ⁻¹	0.007298	6.3394	0.9006	26.62
Zhang <i>et al.</i> (2017), VS4: $h_q = 1.6\text{m}$, vehicle speed = 36 km hr ⁻¹ , Wind 36 km hr ⁻¹	0.005411	5.6387	0.9339	30.67
Zhang <i>et al.</i> (2017), VS5: $h_q = 1.6\text{m}$, vehicle speed = 36 km hr ⁻¹ , Wind 54 km hr ⁻¹	0.003478	4.3150	0.8574	37.89

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- Note that only two contour lines were available for retrieving TKE and height values from this vehicle within Figure 14 of Kim *et al.* (2016); while the correlation coefficient is formally unity, this is a two-point line.
- Woodward *et al.* (2019) Figure 21 turbulent velocity components in the parallel and transverse directions were squared, and distances were scaled to give equivalent distances from wind-tunnel to ambient environment.



895 **Table 2.** Model performance for NO₂, PM_{2.5}, and O₃, 10km grid cell size North American domain. No VIT refers to simulation without
 896 vehicle-induced turbulence, VIT refers to the simulation incorporating vehicle-induced turbulence. **Bold-face** print identifies the better
 897 score, italics the worse score, and regular font indicates similar performance, between the two simulations, for each metric and chemical
 898 species compared.

Species	Evaluation Metric	North America		Canada		USA	
		No VIT	VIT	No VIT	VIT	No VIT	VIT
NO ₂ (ppbv)	FAC2	<i>0.449</i>	0.474	<i>0.437</i>	0.464	<i>0.461</i>	0.484
	MB	<i>1.195</i>	0.142	<i>1.553</i>	0.716	<i>0.860</i>	-0.396
	MGE	<i>4.226</i>	3.542	<i>3.679</i>	3.057	<i>4.738</i>	3.996
	NMGE	<i>0.832</i>	0.698	<i>0.911</i>	0.757	<i>0.783</i>	0.661
	r	0.515	<i>0.511</i>	0.520	<i>0.518</i>	0.507	<i>0.506</i>
	RMSE	<i>7.089</i>	5.665	<i>6.058</i>	4.764	<i>7.934</i>	6.396
	COE	<i>-0.083</i>	0.092	<i>-0.238</i>	-0.029	<i>-0.017</i>	0.142
	IOA	<i>0.459</i>	0.546	<i>0.381</i>	0.486	<i>0.492</i>	0.571
PM _{2.5} (µg m ⁻³)	FAC2	<i>0.451</i>	0.453	<i>0.402</i>	0.412	0.466	<i>0.465</i>
	MB	-2.116	<i>-2.619</i>	-0.032	<i>-0.669</i>	-2.688	<i>-3.154</i>
	MGE	<i>4.982</i>	4.733	<i>4.733</i>	4.237	<i>5.043</i>	4.864
	NMGE	<i>0.672</i>	0.638	<i>0.879</i>	0.787	<i>0.632</i>	0.610
	r	<i>0.185</i>	0.211	<i>0.147</i>	0.163	<i>0.217</i>	0.241
	RMSE	<i>7.933</i>	7.300	<i>8.870</i>	7.323	<i>7.628</i>	7.271
	COE	<i>-0.203</i>	-0.143	<i>-0.431</i>	-0.281	<i>-0.188</i>	-0.146
	IOA	<i>0.399</i>	0.429	<i>0.285</i>	0.360	<i>0.406</i>	0.427
O ₃ (ppbv)	FAC2	<i>0.819</i>	0.823	<i>0.760</i>	0.767	<i>0.830</i>	0.833
	MB	<i>-0.097</i>	0.080	<i>-3.652</i>	-3.498	0.503	<i>0.684</i>
	MGE	<i>10.050</i>	10.009	<i>8.111</i>	8.023	<i>10.379</i>	10.346
	NMGE	<i>0.325</i>	0.323	<i>0.343</i>	0.339	<i>0.322</i>	0.321
	r	0.707	0.707	<i>0.703</i>	0.705	0.694	0.694
	RMSE	<i>13.095</i>	13.035	<i>10.357</i>	10.242	<i>13.511</i>	13.458
	COE	<i>0.239</i>	0.242	<i>0.144</i>	0.153	<i>0.229</i>	0.232
	IOA	<i>0.619</i>	0.621	<i>0.572</i>	0.577	<i>0.615</i>	0.616

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901 **Table 3.** Model performance for NO₂, PM_{2.5}, and O₃, 2.5 km grid cell size Pan Am domain. No VIT refers to simulation without vehicle-
 902 induced turbulence, VIT refers to the simulation incorporating vehicle-induced turbulence. **Bold-face** print identifies the better score, italics
 903 the worse score, and regular font indicates similar performance, between the two simulations, for each metric and chemical species compared.

Species	Evaluation Metric	PanAm Domain July		PanAm Domain January	
		No VIT	VIT	No VIT	VIT
NO ₂ (ppbv)	FAC2	<i>0.584</i>	0.593	0.714	<i>0.711</i>
	MB	<i>1.005</i>	0.386	<i>0.852</i>	-0.328
	MGE	<i>4.137</i>	3.866	<i>5.166</i>	5.146
	NMGE	<i>0.670</i>	0.626	<i>0.457</i>	0.455
	r	0.560	<i>0.543</i>	0.736	<i>0.693</i>
	RMSE	<i>6.909</i>	6.373	<i>7.917</i>	7.892
	COE	<i>0.059</i>	0.121	<i>0.348</i>	0.350
	IOA	<i>0.530</i>	0.560	<i>0.674</i>	0.675
PM _{2.5} (μg m ⁻³)	FAC2	0.573	<i>0.569</i>	<i>0.563</i>	0.592
	MB	-2.669	<i>-3.055</i>	<i>3.930</i>	2.362
	MGE	<i>5.813</i>	5.729	<i>8.315</i>	7.012
	NMGE	<i>0.537</i>	0.529	<i>0.865</i>	0.729
	r	<i>0.338</i>	0.346	<i>0.163</i>	0.170
	RMSE	<i>8.972</i>	8.791	<i>24.875</i>	23.194
	COE	<i>-0.077</i>	-0.061	<i>-0.463</i>	-0.234
	IOA	<i>0.462</i>	0.467	<i>0.269</i>	0.383
O ₃ (ppbv)	FAC2	<i>0.831</i>	0.832	<i>0.852</i>	0.854
	MB	4.138	<i>4.213</i>	1.652	<i>1.731</i>
	MGE	10.640	<i>10.648</i>	<i>6.433</i>	6.427
	NMGE	0.333	0.333	0.259	0.259
	r	0.709	0.709	0.688	<i>0.687</i>
	RMSE	13.826	<i>13.838</i>	<i>8.440</i>	8.427
	COE	0.146	0.146	<i>0.190</i>	0.191
	IOA	0.573	0.573	<i>0.595</i>	0.596

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