



1 The dynamic-thermal structures of the planetary boundary layer dominated by
2 synoptic circulations and the regular effect on air pollution in Beijing

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17 **Abstract.** Synoptic circulations play important roles in meteorological conditions and air quality within the
18 planetary boundary layer (PBL). Based on Lamb-Jenkinson weather typing and multiple field measurements, this
19 study reveals the mechanism of how the coupling effects of multiscale circulations influence PBL structure and
20 pollution. Due to the topographic blocking in the daytime, pollutants accumulate in the plain areas horizontally.
21 The sinking divergent flows overlying on the rising convergent flows within the PBL inhibit the continuously
22 upward dispersion of aerosols vertically. At night, the horizontal and vertical coupling mechanisms synergistically
23 worsen the pollution. The large-scale environmental winds and regional-scale breezes affect the pollution directly
24 via the horizontal coupling effect, which generates a pollution convergent zone of different directional flows. The
25 relative strength of flows causes the severely polluted area to move around horizontally from 39°N to 41°N. In
26 addition, the multiscale circulations regulate the mixing and diffusion of pollutants indirectly via the vertical
27 coupling effect, which changes the PBL dynamic-thermal structure. The warm advection transported by the upper
28 environmental winds overlies the cold advection transported by the lower regional breezes, generating strong
29 wind direction shear and advective inversion. The capping inversion and the convergent sinking motion within the
30 PBL suppress massive pollutants below the zero speed zone. The multilayer PBL under cyclonic circulation has no
31 diurnal variation. Weak ambient winds strengthen the mountain breezes observably at night, the temperature
32 inversion can reach 900 m. The nocturnal shallower PBL, consistent with the zero velocity zone between ambient
33 and mountain winds, can reach 600 m. By contrast, the PBL under southwesterly circulation is a mono-layer with
34 obvious diurnal variation, reaching 2000 m in the daytime. The strong winds circulations restrain the
35 development of regional breezes, the zero speed zone is located at 400 m and the inversion is lower than 200 m
36 at night. The PBL under westerly circulation has a hybrid structure with both multiple aerosol layers and diurnal
37 variation. The inversion is generated by the vertical shear of zonal winds. Clean and strong north winds are
38 dominated under anticyclone circulation, the vertical shear and the diurnal variation of thermal field disappear
39 because of strong turbulent mixing, and there is no significant PBL structure. Our results imply that the algorithm
40 of atmospheric environmental capacity under synoptic circulations, such as the cyclonic type, with a multilayer
41 PBL needs to be improved.



42 **Keywords** Synoptic Circulation Types, Planetary Boundary Layer, Multiscale Circulations Coupling, Regional
43 Breezes, Air Pollution

44 1. Introduction

45 Beijing is the political, economic and cultural center of China. With the recent economic development and
46 acceleration of urbanization, an increasing number of air pollution episodes have emerged and pose a direct
47 threat to human health (Quan et al., 2014; Fu et al., 2014; Cheng et al., 2016; Song et al., 2017). Thus, numerous
48 comprehensive observations and studies on the planetary boundary layer (PBL) and air pollution have been
49 carried out in recent years. Severe pollution is closely related to emissions (Zhang et al. 2012; Wang and Chen
50 2016), synoptic circulations (Wang et al., 2014; Wu et al., 2017; Liao et al., 2017; Miao et al., 2017a, b),
51 topography (Wang et al., 2018; Zhang et al., 2018) and physical and chemical reaction processes (Sun et al., 2015;
52 Zheng et al., 2015a; Yang et al., 2016). In addition to local emissions in Beijing, massive pollutants are generated
53 in southern Hebei Province and transported northward to Beijing through regional transportation (Miao et al.,
54 2016; Chang et al., 2018; Han et al., 2018). Emissions in a particular area normally do not change much over a
55 short period; however, large-scale atmospheric circulations play a leading role in the transportation, accumulation
56 and dispersion of pollution and thus result in the day-to-day variation of air pollutants (Tai et al. 2012; Zhang 2017;
57 Wang et al., 2018). Zheng et al. (2015b) explored the relationships between AOD and synoptic circulations and
58 found that a uniform surface pressure field in eastern China or a steady straight westerly in the middle
59 troposphere is typically responsible for heavy pollution events. Miao et al. (2017a) specially targeted summertime
60 synoptic types, indicating that the horizontal transport of pollutants induced by the synoptic forcing is the most
61 important factor affecting the air quality of Beijing in summer. They also found that synoptic patterns with
62 high-pressure systems located to the east or southeast of Beijing are the most favorable types for heavy aerosol
63 pollution events. Li et al. (2020) quantitatively analyzed the contributions of different large-scale circulations on
64 PM_{2.5}.

65 Beijing is located in the North China Plain (NCP) and is surrounded by Yan and Taihang Mountains to the
66 north and west, respectively (Fig. 1b). This semibasin topography blocks and decelerates the relatively weak
67 southerly flows (Li et al., 2007). Aerosol pollutants from southern provinces through regional transportation
68 stagnate and converge in front of the mountains, leading to the accumulation zone of pollution. In addition, Bohai
69 Sea lies to the southeast and is approximately 150 km from Beijing. This unique geographic location and
70 topography results in diurnal variations in the mountain-plain breeze (MPB) and sea-land breeze (SLB) under
71 relative weak synoptic flows. The SLB can penetrate deep into the mainland when it is blooming, and aerosol
72 pollution transported previously over the sea could be recirculated to the Beijing-Tianjin-Hebei region (Liu et al.,
73 2009; Miao et al., 2017a; Bei et al., 2018). As Beijing is surrounded by mountains and relatively far from the Bohai
74 Sea, the intensity of the MPB circulation is much stronger compared to the sea-land breeze circulation in Beijing
75 (Chen et al., 2009; Miao et al., 2015a, b), especially when synoptic circulations dominate in Bohai areas. Miao et
76 al. (2015b) found that the regional-scale MPB circulations can modulate aerosol pollution by lifting or suppressing
77 PBL. Chen et al. (2009) found that the MPB played an important role in the vertical transportation and dispersion
78 of pollutants via the mountain chimney effect.

79 The PBL structure is also a key factor affecting the distribution and intensity of pollutants in addition to the
80 circulations. The thermal structure of the PBL determines the vertical dispersion of aerosols. In the daytime
81 convective layer, air pollution tends to be mixed vertically and homogeneously because of intensified turbulence
82 and eddies of different sizes by radiation (Stull, 1988). After sunset, the turbulence decays and a stable boundary
83 layer forms with weak turbulence. A radiation inversion on the ground caps the pollutants and leads to the
84 accumulation near the surface. Hu et al. (2014) found that westerly warm advection above the Loess Plateau was



85 transported over the NCP and imposed a thermal inversion, which acted as a lid and capped the pollution in the
86 boundary layer. The dynamic structure of the PBL, including wind shears and turbulence, can modify air quality by
87 influencing the dispersion and transport processes of air pollutants (Li et al., 2019). Zhang et al. (2020) found that
88 a much weaker vertical wind shear was observed in the lower part of the PBL under polluted conditions,
89 compared with that under clean conditions, which could be caused by the strong ground-level PM_{2.5}
90 accumulation induced by weak vertical mixing in the PBL. However, due to the lack of comprehensive observation
91 with high vertical resolution, the dynamic and thermal PBL structure, as well as the mechanisms of how the
92 synoptic circulations and regional-scale circulations influence the PBL structure and air quality, is not well
93 understood. Therefore, the relationships among the multiscale circulations, PBL structure and air pollution should
94 be studied in depth. Many classification approaches have been used to discuss the distinctions of different
95 synoptic circulations, which can be mainly divided into subjective and objective methods. Objective weather
96 classification methods have the advantages of convenient operation, high objectivity and efficiency, hence they
97 have been employed widely in recent years (Zhang et al., 2016; Ye et al., 2016; Miao et al., 2017a). In this study,
98 we adopt an objective Lamb-Jenkinson classification scheme to categorize the large-scale atmospheric
99 circulations centered on Beijing. The Lamb-Jenkinson approach has been applied in many previous studies (Huang
100 et al., 2016; Liao et al., 2017; Yu et al., 2017), which have confirmed that the categorization results have clear
101 physical understanding.

102 This study is based on different synoptic circulations and attempts to investigate the synergetic effects of
103 multiscale circulations on the PBL dynamic-thermal structure and air pollution in detail. The remainder of this
104 paper is organized as follows. Sect. 2 describes the instruments, data and method. Sect. 3 classifies the synoptic
105 circulation types and selects typical types as research objects. Moreover, it further investigates how the coupling
106 mechanism of synoptic circulations and regional-scale circulations changes the dynamic and thermal PBL
107 structure and air pollution. Sect. 4 summarizes the primary conclusions.

108 2. Data and Method

109 2.1 Meteorological data

110 The daily mean sea level pressure (MSLP) and wind fields at 850 hPa were obtained from the National
111 Center for Atmospheric Research (NCAR) reanalysis data (gridded at $2.5^\circ \times 2.5^\circ$). The divergence and vertical
112 velocity reanalysis data, with a horizontal resolution of $1^\circ \times 1^\circ$ and a temporal resolution of 1 h, were obtained
113 from Re-analysis Interim (ERA-Interim) of European Centre for Medium-Range Weather Forecasts (ECMWF). The
114 hourly mean wind fields at the surface in the Beijing-Tianjin-Hebei area were collected by hundreds of automatic
115 observation data provided by the China Meteorological Administration (CMA).

116 2.2 Remote sensing data

117 The high temporal and spatial resolution data of meteorological fields in the boundary layer are obtained by
118 multiple remote sensing devices. The measuring location of ceilometer, Doppler Lidar and microwave radiometer
119 (MWR) is 39.6°N and 116.2°E , in the courtyard of the Institute of Atmospheric Physics, Chinese Academy of
120 Sciences (Fig. 1b). Steyn et al. (1999) had shown that the aerosol concentration in mixing layer (ML) is close to
121 constant and significantly larger than that in the air above. Thus, the ceilometer (CL51, Vaisala) derives the PBL
122 height by BL-VIEW software according to the minimum value of the local backscatter gradient (Tang et al., 2015).
123 The vertical resolution of the backscatter is 10 meters and the maximal detection range can reach 7.7 km. Three
124 possible PBL heights, with a temporal resolution of 10 minutes, can be output simultaneously to characterize the
125 multiple aerosol layers structure according to the first three largest negative gradients of backscatter. The
126 intensity of backscatter are primarily determined by the concentrations of aerosol particulates; hence, the PBL
127 height derived from the BL-VIEW is a material PBL. A Windcube 100S scanning Doppler Lidar measures the



128 Doppler shift of aerosol particulate backscatter using the light detection and ranging (Lidar) technique. The
129 vertical measuring range is from 50 m to 3.3 km. Several scanning modes are available and the DBS (Doppler
130 Beam Swinging technique) mode, which includes four LOS (lines of sight) spaced 90° apart with a fixed elevation
131 angle and one vertical LOS, is selected to detect the profiles of winds. The vertical resolution of the profiles is 25
132 m and the temporal resolution is 20 s. The temperature and relative humidity profiles in RPG-HATPRO MWR are
133 determined by neural network (NN) algorithm, and the vertical resolution of the profiles is 10–30 m in the lowest
134 0.5 km, 40–90 m from 0.5 km to 2.5 km, 100–200 m from 2 km to 10 km, and the temporal resolution is 1 s.

135 2.3 Pollutant data

136 The hourly PM_{2.5} concentrations in the Beijing-Tianjin-Hebei monitoring sites are acquired from the
137 National Urban Air Quality Real-time Publishing Platform (<http://106.37.208.233:20035/>) issued by the Ministry of
138 Ecology and Environment. Beijing has 35 air quality monitoring stations and other areas have 68 monitoring sites
139 in total.

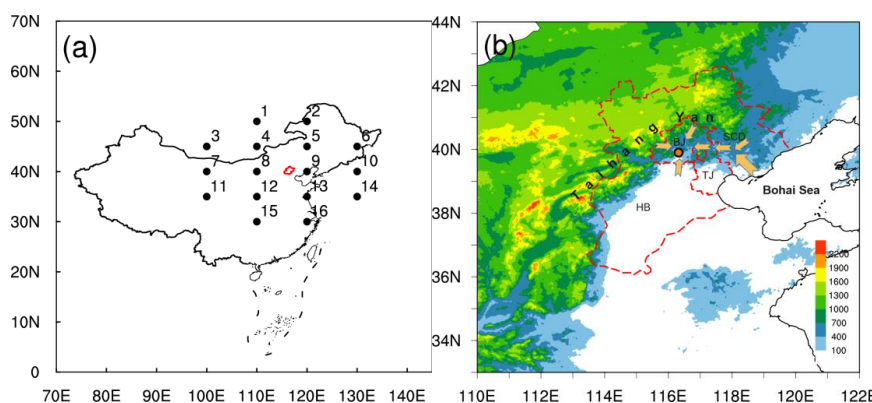
140 2.4 Method

141 The Lamb-Jenkinson weather typing (LWT) approach is widely adopted in large-scale circulation
142 classification (Lamb 1972; Jenkinson and Collison, 1977) because of its automation and explicit meteorologically
143 meaning. To classify the synoptic circulation types, the daily mean sea level pressures (MSLP) in 2018 and 2019
144 were used. The LWT scheme is a half-objective categorization method. The weather patterns are predefined and
145 each day can be identified objectively as one certain type according to a small number of empirical rules (Trigo
146 and DaCamara, 2000). As shown in Fig. 1a, 16 gridded pressure data surrounding the study area (Beijing city)
147 were selected to calculate the direction and vorticity of geostrophic wind. The synoptic circulation can be
148 classified into 26 types in total including two vorticity types (cyclonic, C; anticyclonic, A), eight directional types
149 (northeasterly, NE; easterly, E; southeasterly, SE; southerly, S; southwesterly, SW; westerly, W; northwesterly, NW;
150 and northerly, N), and sixteen hybrid types (CN, CNE, CE, CSE, CS, CSW, CW, CNW, AN, ANE, AE, ASE, AS, ASW, AW,
151 and ANW).

152 The gradient Richardson number (Ri) is the ratio of the buoyancy term to the shear term in the turbulent
153 kinetic equation. It is able to estimate the atmospheric turbulent stability and can be calculated by Eq. 1, where g
154 is the acceleration of gravity and Δz is the height interval between adjacent layers. $\bar{\theta}$ is the mean virtual
155 potential temperature, $\Delta \bar{u}$ and $\Delta \bar{v}$ is the mean zonal and meridional wind speeds within the height interval
156 respectively. Previous studies (Stull, 1988; Guo et al., 2016) suggested that when Ri is smaller than the critical
157 value (0.25), the laminar flow becomes unstable. Thus, we adopt the value of 0.25 as a criterion to determine
158 whether the layer is stable or not.

159

$$Ri = \frac{\frac{g}{\bar{\theta}} \frac{\Delta \bar{\theta}}{\Delta z}}{\left(\frac{\Delta \bar{u}}{\Delta z}\right)^2 + \left(\frac{\Delta \bar{v}}{\Delta z}\right)^2} \quad (1)$$



160
161 Fig. 1 Location of Beijing city in China (red lines). The 16 black points show the location of the $5^{\circ} \times 10^{\circ}$ MSLP grids
162 used for Lamb-Jenkinson weather type classification (a). The terrain height of the North China Plain (shaded, units:
163 m). The filled dots show the locations of remote sensing devices (b). The arrows indicate the horizontal coupling
164 mechanism of how multiscale circulations affect pollution by generating convergent zone.

165 3. Results and Discussions

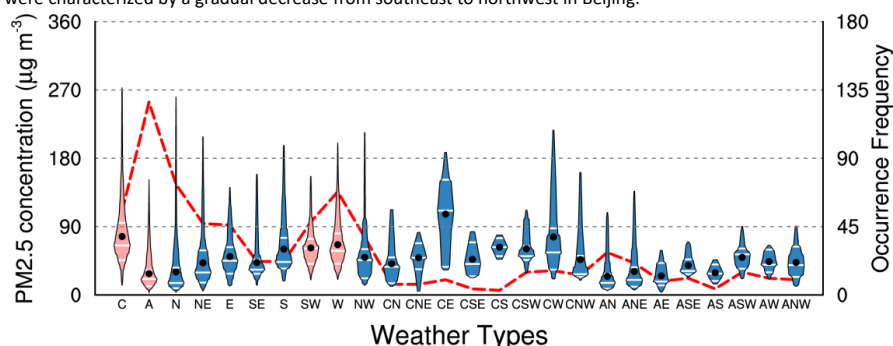
166 3.1 The typical weather types and PM_{2.5} distribution

167 Based on the Lamb-Jenkinson weather typing approach, synoptic circulations from 2018 to 2019 were
168 classified into predefined 26 circulation patterns and each day has a specific type. The distributional
169 characteristics of daily averaged PM_{2.5} concentration, as well as the occurrence frequency of different circulation
170 patterns, were statistically conducted. The occurrence frequencies of the two vorticity and eight directional types
171 were much higher than those of the other sixteen hybrid types, accounting for 75% of total days (Fig. 2).
172 According to the pollution intensity, three pollution types (cyclonic C, southwesterly SW and westerly W) and one
173 clean type (anticyclonic A) occurring most frequently in the NCP were selected as the studied circulation patterns.
174 It was consistent with the results of Li et al. (2020) on the relationship between pollutant concentration and
175 circulation types in northern China. Weather types with high PM_{2.5} concentration but occurring no more than
176 ten times, such as type CE and type CW, were not discussed in this article. The average and extreme PM_{2.5}
177 concentrations of type C reached $77 \mu\text{g}/\text{m}^3$ and $270 \mu\text{g}/\text{m}^3$, respectively, and were much stronger than the other
178 pollution types. Clearly, the cyclonic circulation pattern was more conducive to severe pollution events. The
179 circulation of type A was the most common type, and the PM_{2.5} concentration was $28 \mu\text{g}/\text{m}^3$, which was the
180 lowest.

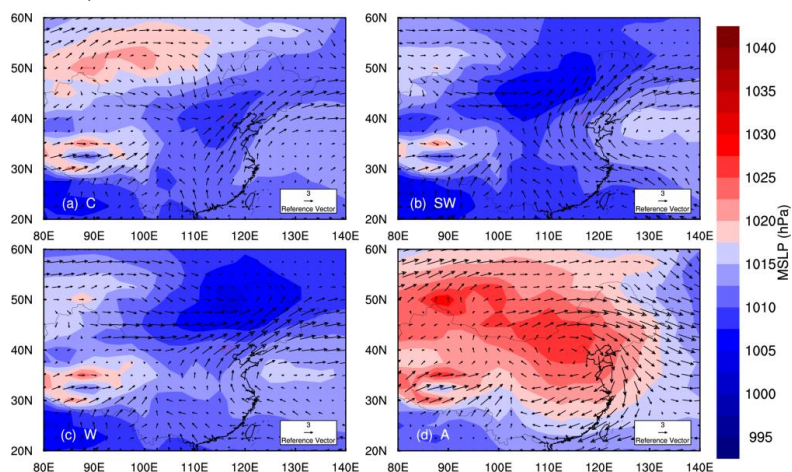
181 As shown in Fig. 3, the locations of the high and low pressures and the intensity of the wind fields at 925
182 hPa under different circulation patterns were clearly distinct. In type C, Beijing was located in the center of low
183 pressure, and the sea to the east of China was controlled by an anticyclone (Fig. 3a). Southwesterly winds
184 prevailed, flowing northward to Beijing along the periphery of the anticyclone with an average wind speed of 3
185 m/s. In type SW, Beijing lay southeast of the low pressure in Mongolia, and the high pressure over the sea was
186 significantly enhanced compared with type C (Fig. 3b). Therefore, southeasterly winds prevailed to the south of
187 Beijing and shifted southwesterly after flowing by. In type W, westerly winds were dominant and converged with
188 southwesterly flows to the north of Beijing (Fig. 3c). The mean velocity of environmental flows in type SW and
189 type W was observably larger than that in type C. In general, the mainland was mainly controlled by low pressure
190 with an anticyclone lying over the sea to the east of China in pollution types C, SW and W, and southerly flows
191 dominated at 925 hPa. By contrast, northern China in the clean type A was occupied by high pressure. Beijing was



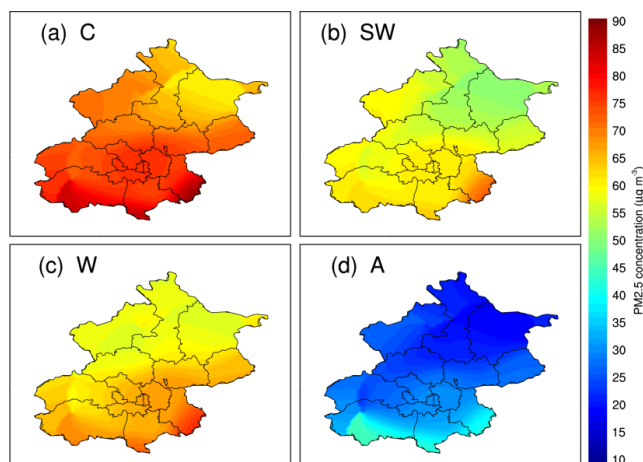
192 located in the center of high pressure with strong northerly winds in the lower level (Fig. 3d).
193 The pollution intensity is closely related to the large-scale weather circulations. Although the dominant
194 synoptic patterns in different seasons vary greatly, the modulating effects on air pollution of specific circulation
195 types in different seasons are similar (Liao et al., 2017; Li et al., 2020). The spatial distribution of PM_{2.5} in Beijing
196 under pollution types C, SW, W and clean type A is shown in Fig. 4. Type C had the highest pollution level, with
197 the PM_{2.5} concentration increasing from 60 $\mu\text{g}/\text{m}^3$ in the northwestern mountainous area to 90 $\mu\text{g}/\text{m}^3$ in the
198 south-central plain area, which was significantly higher than the values for types SW and W. Type A was highly
199 ventilated, with a PM_{2.5} concentration below 30 $\mu\text{g}/\text{m}^3$ in most areas. Under the influence of semibasin
200 topography surrounded by mountains on three sides (Fig. 1b), the pollution concentrations in all weather types
201 were characterized by a gradual decrease from southeast to northwest in Beijing.



202
203 Fig. 2 Daily averaged PM_{2.5} concentration (box plots, units: $10^{-1} \mu\text{g m}^{-3}$) and the occurrence frequencies of 26
204 weather types (red dashed lines) from 2018 to 2019. The red boxes represent classical types selected for research.
205 The black dots represent the mean values.



206
207 Fig. 3 The daily MSLP (shaded, units: hPa) and wind fields at 925 hPa (vectors, units: m s^{-1}) for types C, SW, W and
208 A from 2018 to 2019.



209
210 Fig. 4 The averaged PM_{2.5} concentration (shaded, units: $10^{-1} \mu\text{g m}^{-3}$) in Beijing for types C, SW, W and A from
211 2018 to 2019.

212 3.2 The flow field and dynamic-thermal structure of the PBL under typical weather 213 types

214 As mentioned above, due to the special topography and geographical location in Beijing, both large-scale
215 weather circulations and regional-scale thermal circulations have conspicuous effects on modulating pollution. In
216 addition, the thermal and dynamic structure of the PBL also has an appreciable impact on the mixing and
217 diffusion of pollutants. Therefore, the multiscale circulations can not only influence the pollution directly but also
218 influence it by changing the PBL structure indirectly. To reveal the mechanisms of how the coupling effects of
219 multiscale circulations affect the PBL structure and air pollution under different synoptic patterns, we conduct an
220 analysis of the horizontal flow field and vertical PBL structure in depth by choosing typical cases lasting two days
221 in the same weather type (C, SW, W and A).

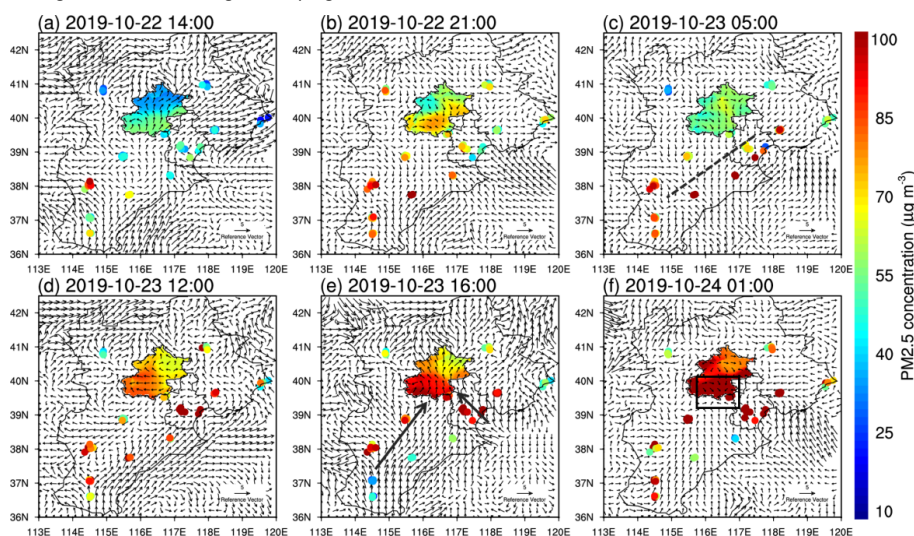
222 3.2.1 Multilayer PBL structure under type C circulation

223 The mainland was governed by low pressure under type C synoptic circulations, and the ambient winds
224 were mainly southwesterly (Fig. 3a). On the afternoon of the 22nd, the plain breezes in central Hebei, which were
225 induced by thermal contrast between the mountain and plain, blocked weak environmental winds and the direct
226 transportation of pollutants to Beijing (Fig. 5a). The westerly and the northerly mountain breezes began to prevail
227 at night while the conversion from sea breeze to land breeze was not obvious (Fig. 5b). The onshore winds in the
228 coastal area were notably larger than the northerly mountain breezes in southern Chengde (SCD), which were
229 diverted to the west and east. The diverted easterly winds converged with the onshore winds, enhancing the
230 easterly winds and the east pollution transport channel. Sun et al. (2019) have found that the pressure gradients
231 between the plain and mountain areas are critical causes of the easterly winds in Beijing. Consequently, easterly
232 winds gathered with mountain breezes and formed a pollution convergent zone. Weak environmental winds not
233 only made the pollution channels hard to establish but also caused the pollutants to recirculate southward by
234 strong downslope breezes further in the early morning (Fig. 5c). A mesoscale convergent belt was generated in
235 southeastern Hebei, providing conditions for the transportation of pollutants later. At noon on the 23rd, the
236 intensified plain winds transported high concentrations of aerosols from the right side of the convergent belt to
237 Beijing (Fig. 5d). Large-scale environmental winds were strengthened and dominated in the afternoon (Fig. 5e),
238 leading to the establishment of the south and east pollution transport channels and further exacerbating the air
239 quality. On the night of the 23rd, easterly winds were observably strengthened again, joining with the downslope



240 breezes and the ambient southerly flows (Fig. 5f). The four directional airflows formed a convergent zone that
241 caused pollutants to accumulate dramatically in the plain areas. This convergent region that is generated by the
242 coupling effect of large-scale circulation and regional-scale mountain breezes at night also appeared in other
243 pollution types, as will be discussed later.

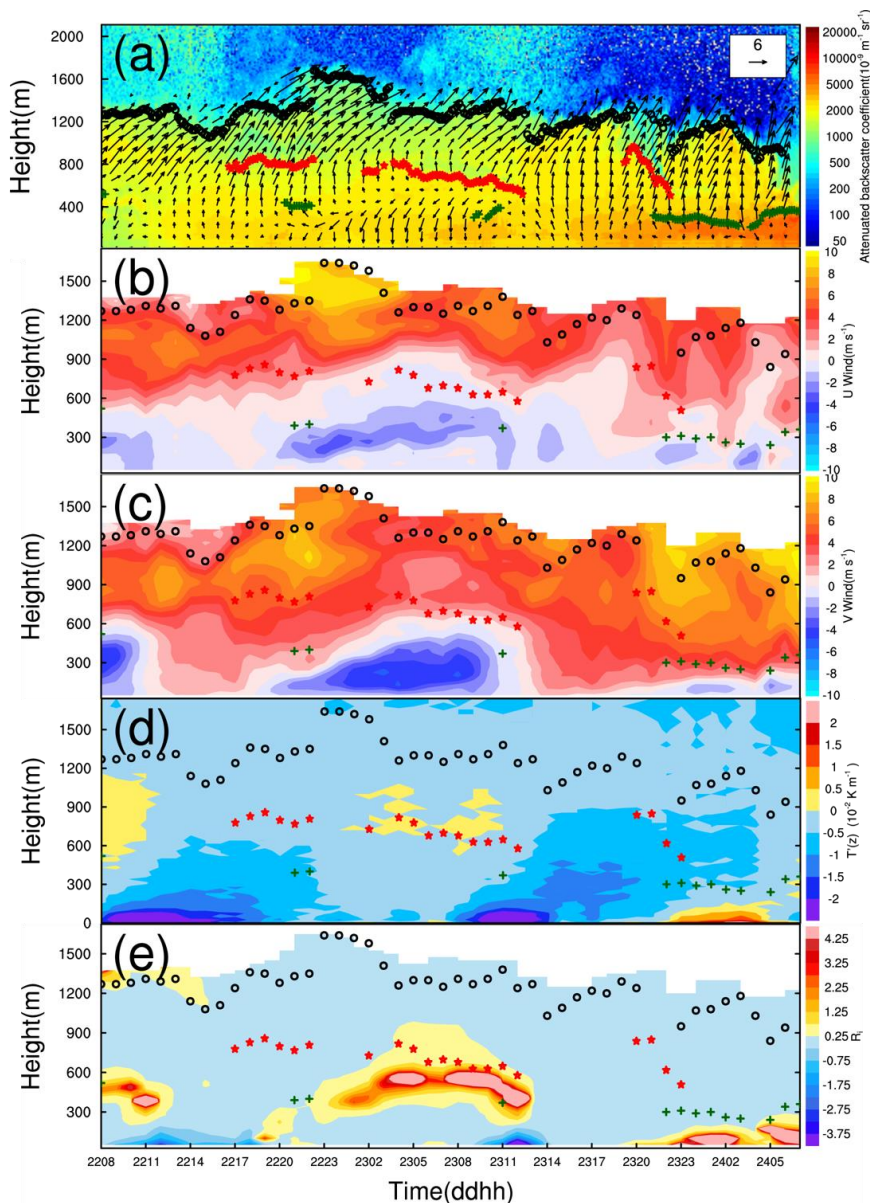
244 The PBL under type C circulation presented a multilayer structure without diurnal variation (Fig. 6a). The
245 highly stable structure and weak ambient winds resulted in a higher aerosol concentration near the surface than
246 that in the other pollution types (Fig. 4). The pollution decreased from bottom to top within the PBL and was
247 characterized by a gradient distribution. It is consistent with previous research (Jiang et al., 2020) that the top PBL
248 height is equal to the maximum detection range of wind Lidar. In the daytime, environmental southwesterly
249 winds dominated within the PBL. On the night of the 22nd, meridional winds turned to easterly (Fig. 5b, 6b), and
250 the northerly downslope winds were strengthened simultaneously in the lower PBL (Fig. 5c, 6c). Easterly and
251 northerly winds were up to 600 m so that the directional shear of meridional and zonal winds ascended
252 considerably. The shallower nocturnal PBL coincided with the zero speed zone between the upper environmental
253 winds and lower regional-scale breezes with the largest directional shear (Fig. 6b, c). Variations of the vertical
254 dynamic structure in the PBL drove the thermal structure to adjust. Warm air advected by large-scale
255 southwesterly winds overlay on the cold air advected by regional-scale northeasterly breezes. Consequently, a
256 conspicuous advective temperature inversion occurred from 600 to 900 m (Fig. 6d) accompanied by stable
257 stratification (Fig. 6e). However, the relatively stronger northerly breezes compared to the environmental winds
258 made the pollutants recirculate southward horizontally (Fig. 5c, 6c). Furthermore, the wind shear developed so
259 high that the stable stratification was above 300 m and the inversion was above 600 m; the pollutants dispersed
260 vertically to some extent. Compared to the previous night, the ambient winds on the night of the 23rd were
261 stronger; thus, both south and east transport channels were established, along with the pollution convergent
262 zone (Fig. 5f). The weak easterly and northerly winds were lower than 300 m (Fig. 6b, c), resulting in temperature
263 inversion and stable stratification connected to the ground. A high concentration of pollution was accumulated in
264 the convergent zone horizontally and trapped below the lowest PBL vertically. Thus, the PM_{2.5} concentration on
265 the night of the 23rd was significantly higher than that on the 22nd.



266
267 Fig. 5 The surface winds (vectors, units: m s^{-1}) in the NCP and PM_{2.5} concentration in Beijing (shaded, units: 10^{-1}
268 $\mu\text{g m}^{-3}$), Hebei and Tianjin (scatter, units: $10^{-1} \mu\text{g m}^{-3}$) for type C. The dashed line represents the convergence belt.



269 The arrow lines represent the pollutant transport channels. The rectangle represents the convergent zone.



270
 271 Fig. 6 Attenuated backscatter coefficient (shaded, units: $10^{-9} \text{ m}^{-1} \text{ sr}^{-1}$) and horizontal winds (vectors, units: m s^{-1})
 272 (a), zonal wind speeds (shaded, units: m s^{-1}) (b), meridional wind speeds (shaded, units: m s^{-1}) (c), gradient of
 273 temperature $T'(z)$ (shaded, units: K km^{-1}) (d), and Richardson number (shaded) (e) for type C. The green crosses,
 274 red stars and black hollow dots represent the lowest, middle and top PBLH, respectively.

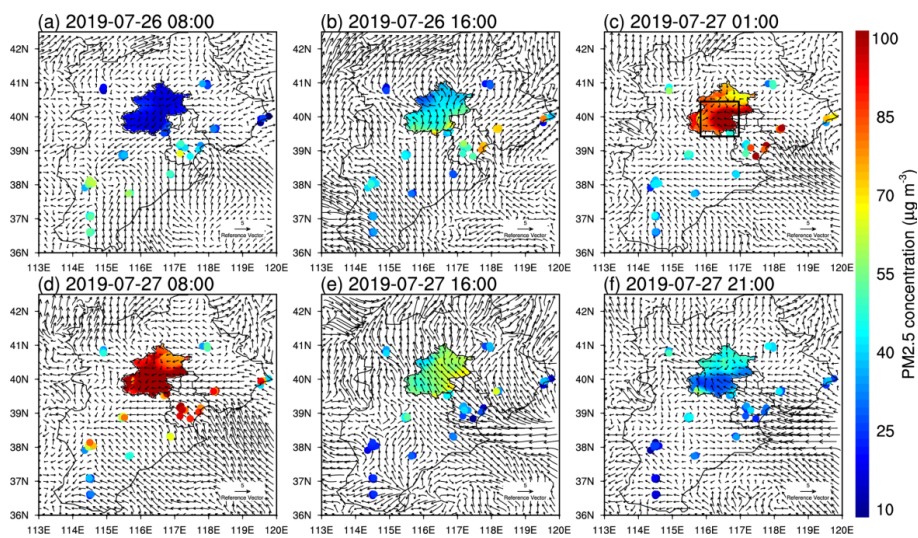
275 3.2.2 Mono-layer PBL structure under type SW circulation

276 Under type SW circulation, the easterly wind component increased in southeastern Hebei and the Bohai Sea,
 277 and the velocity of environmental winds was appreciably higher than that in type C. (Fig. 3b). On the early



278 morning of the 26th, mountain breezes carrying clean air masses prevailed in Beijing, and the air quality was good
279 (Fig. 7a). The basic southerly winds dominated in the Beijing-Tianjin-Hebei region in the afternoon, transporting
280 pollutants northward and causing airflow to converge in plain areas (Fig. 7b). However, pollutants were ventilated
281 horizontally by strong ambient winds and diffused vertically by the intensified turbulent mixing within the
282 growing ML, so the aerosol concentration grew slowly during the day (Fig. 8a). At night, the mountain breezes
283 were strengthened while the ambient southerly winds were weakened; hence, the pollutants were transported to
284 Beijing via the east pollution channel (Fig. 7c). Multiscale circulations of different directions joined and generated
285 a convergent zone in the plain area. Afterwards, easterly flows were further strengthened and transported
286 pollutants to Beijing continuously, the severely polluted area moved westward (Fig. 7d, 8a). In the daytime of the
287 27th, the ambient winds prevailed again, and strong ambient winds removed pollutants by enhancing the
288 ventilation and turbulent mixing (Fig. 7e, 8a). Therefore, the PM_{2.5} concentration decreased instantly and the air
289 quality in the Beijing-Tianjin-Hebei region improved markedly (Fig. 7f).

290 Unlike type C, the PBL presented a monolayer structure in type SW, and the aerosol within the PBL was
291 uniformly distributed (Fig. 8a). Furthermore, the PBL had an obvious diurnal variation and the maximum
292 detection distance of wind Lidar was only consistent with the top ML in type SW. The nocturnal PBL and the
293 growing or collapsing ML were usually lower than the maximum detection distance, indicating that there were
294 residual aerosols above the PBL. In the daytime of the 26th, southwesterly winds dominated within the PBL, and
295 the temperature lapse rate was greater than 0.5 °C/100 m. Along with radiation reinforcing turbulent kinetic
296 energy, the PBL rose to 1200 m. Pollutants were transported to Beijing but mixed vertically (Fig. 8a), so the PM_{2.5}
297 concentration near the surface grew slowly (Fig. 7b). On the night of the 26th, the regional-scale circulation
298 developed upward, and the vertical wind shears between the lower regional breezes and upper environmental
299 winds were strengthened prominently (Fig. 8b, c). The warm advection overlay on the cold advection resulted in
300 advective inversion, forcing the PBL to adjust to become stable, correspondingly (Fig. 8d, e). Similar to type C, a
301 high concentration of pollutants was trapped below the zero wind speed zone where the nocturnal PBL was
302 located. In the daytime of the 27th, large-scale environmental winds within the PBL were strengthened greatly.
303 The PBL height was 800 m higher than that of the previous day; thus, the pollutants were advected horizontally
304 and diffused vertically (Fig. 8a). The basic southerly winds with high speed prevailed in central and southern
305 Beijing on the night of the 27th, preventing the mountain winds from flowing southward (Fig. 7f). As a result, no
306 vertical shear of meridional winds occurred in the dynamic field (Fig. 8c) and no temperature inversion occurred
307 in the thermal field (Fig. 8d). The PM_{2.5} concentration was further reduced. It can be inferred that the
308 temperature inversion in type SW was generated by the vertical thermal contrast of meridional winds. When the
309 meridional winds were uniformly southerly winds within and above the PBL, the air masses in the upper layer had
310 the same thermal properties as that in the lower layer, which will reduce the vertical wind shear and destroy the
311 stable inversion structure.

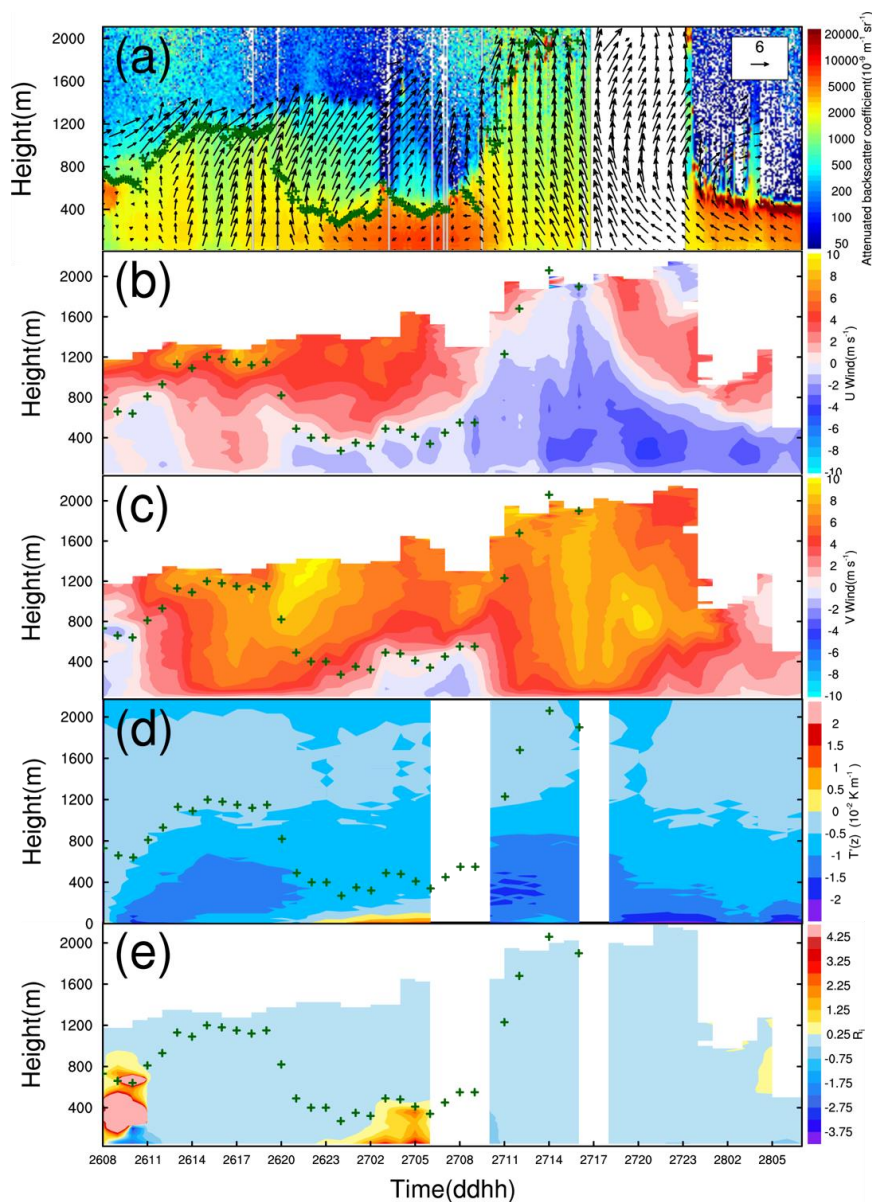


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Fig. 7 The surface winds (vectors, units: m s^{-1}) in the NCP and PM2.5 concentration in Beijing (shaded, units: $10^{-1} \mu\text{g m}^{-3}$), Hebei and Tianjin (scatter, units: $10^{-1} \mu\text{g m}^{-3}$) for type SW. The rectangle represents the convergent zone.



315

316 Fig. 8 Attenuated backscatter coefficient (shaded, units: $10^{-9} \text{ m}^{-1} \text{ sr}^{-1}$) and horizontal winds (vectors, units: m s^{-1})
317 (a), zonal wind speeds (shaded, units: m s^{-1}) (b), meridional wind speeds (shaded, units: m s^{-1}) (c), gradient of
318 temperature $T'(z)$ (shaded, units: K km^{-1}) (d), and Richardson number (shaded) (e) for type SW. The green crosses
319 represent the PBLH.

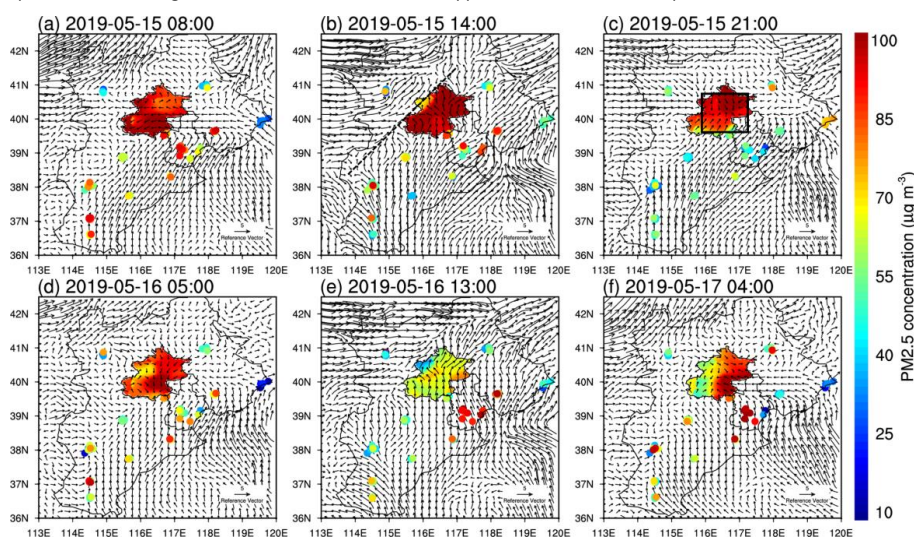
320 3.2.3 Hybrid structure PBL under type W circulation

321 Under type W circulation, strong easterly winds transported a high concentration of aerosols to Beijing
322 through the east pollution channel, and the PM_{2.5} concentration had already reached a high level in the early
323 morning (Fig. 9a). Taking the mountain as the boundary, environmental westerly winds prevailed in northwestern

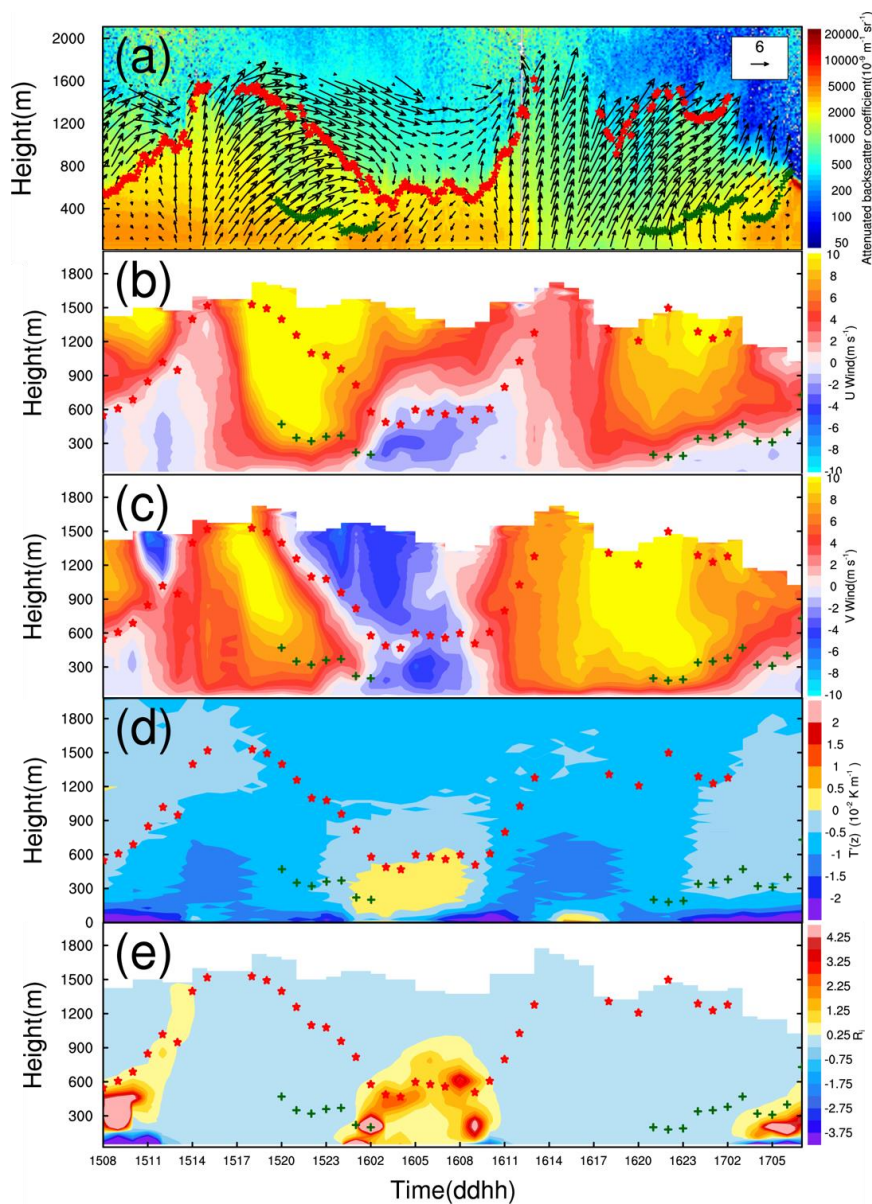


324 Hebei and southwesterly winds prevailed in southern Hebei in the afternoon. The two directional flows carried
325 pollutants and formed a convergent belt along the western mountains (Fig. 3c, 9b). This distribution of synoptic
326 circulations in type W was conducive to the occurrence of severe pollution around mountains. Similar to other
327 pollution types, the ambient winds converged with region-scale mountain breezes at night, forming a convergent
328 zone (Fig. 9c). The convergent zone moved southward later because of intensified mountain breezes (Fig. 9d). The
329 large velocity of environmental winds leads to strong ventilation (Fig. 9e). In addition, the increasing PBL made
330 the pollutants diluted vertically, and the air pollution was alleviated temporarily. On night of the 16th (Fig. 9f), the
331 synergistic effects of multiscale circulations led to the convergent zone again, and pollution occurred in the
332 easterly flows with a high PM2.5 concentration.

333 The PBL under type W circulation presented a hybrid structure, having similar characteristics of types C and
334 SW simultaneously. Similar to type C, the aerosol concentration was characterized by a gradient distribution
335 within the multilayer PBL (Fig. 10a). However, the PBL had an obvious diurnal variation, and the maximum
336 detection distance of wind Lidar was only consistent with the top ML in the daytime, similar to type SW. Although
337 the PBL height reached 1600 m in the daytime (Fig. 10a), the PM2.5 concentration at the surface did not decrease
338 observably because of the massive pollution accumulated previously and the continuous emissions and
339 transportation of pollutants (Fig. 9b). The mixing layer collapsed along the zero wind speed of meridional winds
340 after sunset, and the breezes within nocturnal PBL shifted northwesterly at night (Fig. 10b, c). In type W, zonal
341 circulation dominated. The vertical shear of zonal winds was intensified significantly at night, while the vertical
342 shear of meridional winds diminished. Therefore, it can be assumed that the temperature inversion in type W was
343 produced by the vertical shear of zonal winds. The thermal contrast between the upper westerly winds and the
344 lower easterly winds produced a deep inversion layer that existed from the surface to 500 m (Fig. 10d), as well as
345 a stable stratification with a depth exceeding 600 m (Fig. 10e). This is consistent with the findings of Hu et al.
346 (2014) that westerly warm advection from the Loess Plateau was transported over the NCP and imposed a
347 thermal inversion above the PBL. The top of the PBL was consistent with the top of the inversion and zero wind
348 speed zone, and a high concentration of aerosols was trapped below the zero wind speed zone.



349 Fig. 9 The surface winds (vectors, units: m s^{-1}) in the NCP and PM2.5 concentration in Beijing (shaded, units: 10^{-1}
350 $\mu\text{g m}^{-3}$), Hebei and Tianjin (scatter, units: $10^{-1} \mu\text{g m}^{-3}$) for type W. The dashed line represents the convergence belt.
351 The rectangle represents the convergent zone.
352



353

354 Fig. 10 Attenuated backscatter coefficient (shaded, units: $10^{-9} \text{ m}^{-1} \text{ sr}^{-1}$) and horizontal winds (vectors, units: m s^{-1})
355 (a), zonal wind speeds (shaded, units: m s^{-1}) (b), meridional wind speeds (shaded, units: m s^{-1}) (c), gradient of
356 temperature $T'(z)$ (shaded, units: K km^{-1}) (d), and Richardson number (shaded) (e) for type W. The green crosses
357 and red stars represent the low and top PBLH, respectively.

358 3.2.4 Strong turbulent PBL structure under clean type A circulation

359 Strikingly different from the circulations of pollution types, the mainland was under high pressure control in
360 the clean type, and northwesterly winds with a high velocity carrying clean air masses moved southward (Fig.
361 11a). Strong winds were favorable for the turbulent mixing and the vertical dispersion of pollutants. In addition,



362 the strong ventilation was beneficial to the horizontal spreading of pollutants. Due to the intense turbulent
363 mixing, the vertical wind shear and the diurnal variation of thermal field disappear, and there is no significant PBL
364 structure (Fig. 11a, b). The lapse rate of temperature was greater than $1\text{ }^{\circ}\text{C}/100\text{ m}$, and R_i was less than 0.25
365 within the PBL (not shown). Although the aerosol concentration of the clean type was far less than that of
366 pollution types, the PBL height was only 500 m at night (Fig. 11a). Sometimes, the PBL in the clean type was even
367 lower than that of pollution types, or extended to 2-3 km swiftly because of the instant upward diffusion of
368 aerosol particulates. Unlike pollution types, the PBL height is inconsistent with the maximum detection range of
369 wind Lidar. Therefore, different circulation types should be distinguished when analyzing the long-term
370 relationships between the PBL height and pollution concentration. As shown in Fig. 12 c and d, under the
371 governing of high pressure, descending and divergent airflows of the clean type dominated the whole lower and
372 middle parts of the troposphere, and the sinking velocity was significantly higher than that of pollution types. The
373 vertical velocity changed little vertically due to the northerly winds with a large speed penetrating downward. The
374 intensity of sinking and divergence was higher at night than that in the day, with the strongest divergence
375 occurring near the surface.

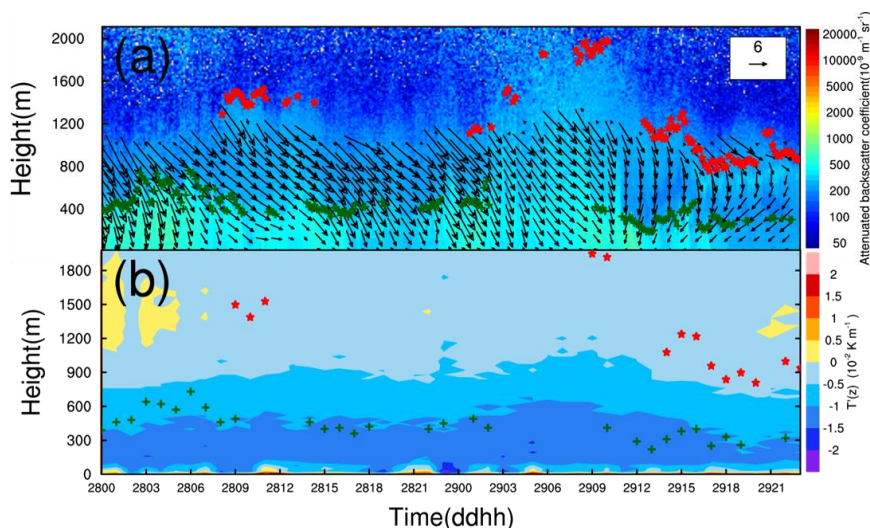
376 3.3 Multiscale circulations coupling mechanism for air pollution

377 In addition to horizontal circulations, the vertical motion of basic flows is also a crucial dynamic factor in
378 forming stable stratification during pollution episodes. The pollution types shared similar vertical motion
379 characteristics as shown in Fig. 12. In the daytime, the NCP region was controlled by a rising motion below 900
380 hPa with a sinking motion overlaying it (Fig. 12a). Correspondingly, the basic flows below 900 hPa presented a
381 convergence, while that above 900 hPa presented a divergence (Fig. 12b). Airflows inside the PBL converged and
382 rose, while the sinking and divergent flows superposed above the PBL, preventing the pollutants from moving
383 upward continuously and making it difficult for the aerosol particulates to diffuse beyond. As a consequence, the
384 pollutants accumulate gradually in the daytime because of the common influences of horizontal topographic
385 blocking and vertical upward mixing with the ML rise. At night, the winds presented a consistent sinking motion
386 below 500 hPa with the largest sinking velocity occurring near the surface (Fig. 12a). Wu et al. (2017) found that
387 the descending motion of synoptic circulations contributed to a reduction in the PBLH by compressing the air
388 mass. In general, the airflow of pollution types is always convergent inside the PBL with the strongest
389 convergence occurring at 950 hPa, regardless of whether it is daytime or nighttime. The height of the nocturnal
390 PBL reduced observably and simultaneously with the convergence zone; meanwhile, divergent downdrafts above
391 the PBL make it difficult for pollutants to diffuse upward (Fig. 12b). Thus, massive pollutants were capped near
392 the surface and accumulated rapidly.

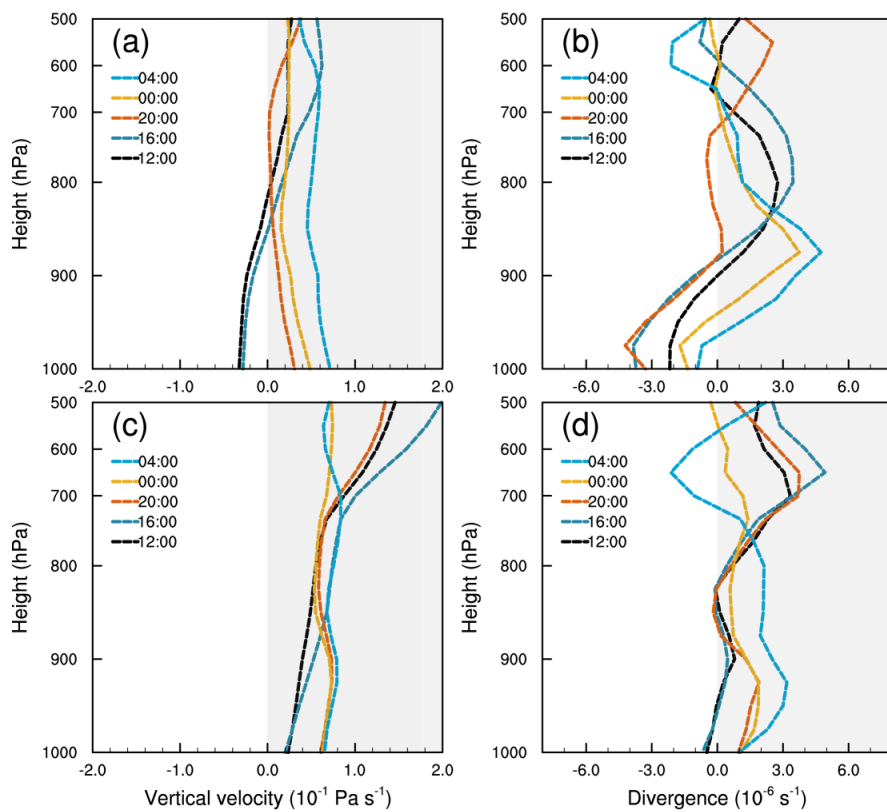
393 To sum up, different pollution patterns (C, SW and W) have similar influential mechanisms that both
394 horizontal and vertical coupling effects of the multiscale circulations have contributed to air pollution. The
395 horizontal coupling mechanism is shown in Fig. 1b. The environmental winds transport pollutants emitted from
396 southern sources to Beijing, mainly through south and east pollution channels. Large-scale environmental winds
397 and regional-scale breezes are coupled, generating a convergent zone of four directional flows horizontally and
398 aggravating the air pollution directly. The relative strength of winds makes the severely polluted area move
399 around horizontally from 39°N to 41°N . The schematic of Fig. 13 demonstrates that the vertical coupling
400 mechanism further influences the mixing and dispersion of pollution indirectly by changing the PBL structure. In
401 the daytime, the sinking divergent flows overlaying the rising convergent flows within the PBL inhibit the
402 continuous upward dispersion of pollutants. At night, the warm advection transported by the upper
403 environmental winds overlies the cold advection transported by the lower regional breezes, generating strong
404 directional wind shear and advective inversion, which are near the top of regional breezes. This dynamic structure
405 forces the PBL to be a stable stratification. The nocturnal PBL is located at the zero speed zone between the



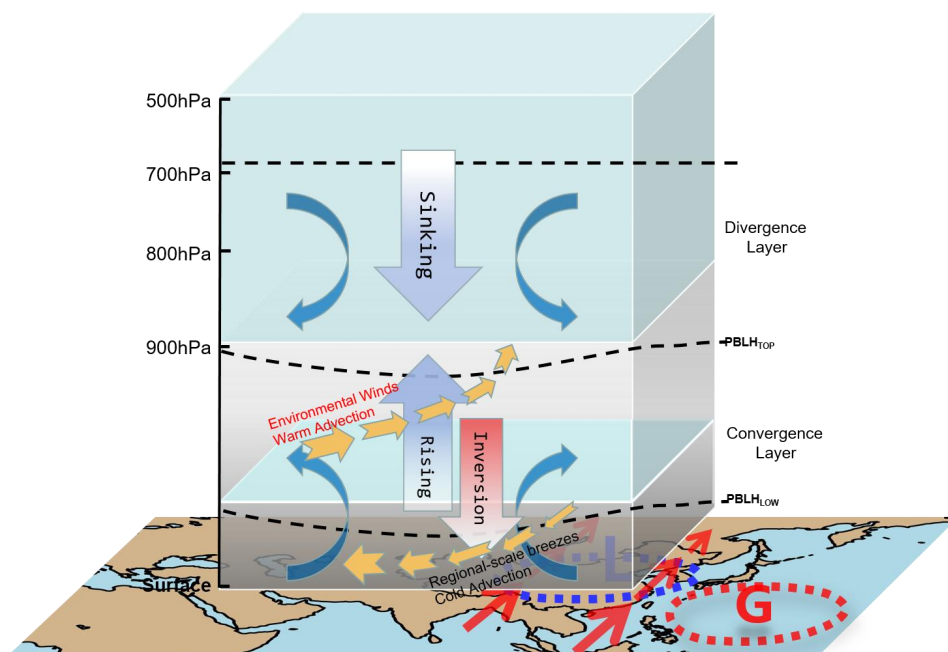
406 regional-scale breezes and the environmental winds, and the relative strength of winds determines the PBL height.
407 The capping inversion cooperating with the convergent sinking motion within the PBL suppresses massive
408 pollutants below the zero speed zone.
409 However, the flow field and the PBL dynamic-thermal structure under different synoptic circulations vary
410 widely with the location and intensity of high and low pressure and wind fields, resulting in differences in
411 pollution. The multilayer PBL under type C circulation has no obvious diurnal variation. Weak ambient winds
412 strengthen the mountain breezes observably at night. Thus, the temperature inversion and zero speed zone can
413 reach 600 m to 900 m vertically, and the pollution convergent zone occurs in the plain areas horizontally. By
414 contrast, the PBL under type SW circulation is a mono-layer with obvious diurnal variation, reaching 2000 m in
415 the daytime. The strong environmental winds restrain the development of regional breezes, the zero speed zone
416 is located at 400 m and the temperature inversion is lower than 200 m at night. The inversion is generated by the
417 vertical shear of meridional winds at night. Southerly winds within and above the PBL having the same thermal
418 properties will diminish the vertical shear and damage the advective inversion structure. The type W circulation is
419 governed by zonal motion and the PBL has a hybrid structure with both multiple aerosol layers and diurnal
420 variations. The vertical comparison of zonal winds leads to a much deeper inversion and stable stratification. The
421 pollution zone under types SW and W circulations is closer to mountainous areas because of strong ambient
422 winds. Furthermore, strong ambient winds make the pollutants ventilate horizontally and diffuse vertically with
423 the growing ML in the daytime.



424
425 Fig. 11 Attenuated backscatter coefficient (shaded, units: $10^{-9} \text{ m}^{-1} \text{ sr}^{-1}$) and horizontal winds (vectors, units: m s^{-1})
426 (a), and gradient of temperature $T'(z)$ (shaded, units: K km^{-1}) (b) for type A. The green crosses and red stars
427 represent the low and top PBLH, respectively.



428
429 Fig. 12 The averaged vertical velocity (units: Pa s^{-1}) (a, c) and divergence (units: 10^{-5} s^{-1}) (b, d) of pollution types (a,
430 b) and the clean type (c, d) in the North China Plain



431
432 Fig. 13 The vertical coupling mechanism of how multiscale circulations affect pollution by changing the PBL
433 dynamic-thermal structure

434 4. Conclusions

435 Based on Lamb-Jenkinson weather typing, the most frequent typical pollution types and clean type were
436 chosen to explore the flow field and the PBL structure under different synoptic patterns. In addition, the
437 horizontal and vertical coupling mechanisms of multiscale circulations, which aggravated pollution synergistically,
438 were further revealed. The results show that different pollution patterns have similar influential mechanisms for
439 air pollution. The large-scale environmental winds and regional-scale breezes, on the one hand, affect the
440 pollution directly via the horizontal coupling effect, which produces a pollution convergent zone. The relative
441 strength of winds makes the severely polluted area move around horizontally. On the other hand, regulate the
442 mixing and diffusion of pollutants indirectly by the vertical coupling effect, which changes the PBL dynamic and
443 thermal structure. Vertical shear between the ambient winds and regional-scale breezes leads to advective
444 inversion and stable stratification, and the relative strength of winds determines the PBL height. Massive
445 pollutants were suppressed below the zero speed zone by the capping inversion and the convergent sinking
446 motion within the PBL. The distinctions of the flow field and PBL dynamic-thermal structure result in the
447 differences of horizontal and vertical pollution, respectively. Based on the fact that both the flow field and PBL
448 structure are dominated by synoptic circulations, the atmospheric environmental capacity (AEC) may vary day by
449 day following the changes in the circulations. Especially when the pollution and meteorological conditions are
450 layered within the PBL, the traditional calculation approach of AEC, which treats the PBL as a uniform and
451 homogenous layer, is no longer applicable. Future work on the quantitative relationships between the PBL
452 structure and air pollution under different weather patterns still needs to be performed. The algorithm of AEC
453 under synoptic circulations with a multilayer PBL, such as cyclonic type circulation, also needs to be improved.

454 Data availability



455 The hourly ground level PM_{2.5} concentration data can be obtained from the National Urban Air Quality
456 Real-time Publishing Platform (<http://106.37.208.233:20035/>). Other data used in this study can be acquired upon
457 request to the corresponding author.

458 Competing interests

459 The authors declare that they have no known competing financial interests or personal relationships that
460 could have appeared to influence the work reported in this paper.

461 Author contribution

462 JY performed the idea, methodology, data processing, visualization and writing. XJ provided writing
463 guidance and funding, revised and polished the paper. WY performed supervision. TG contributed to observation
464 data and discussions of results. ZY provided the research data and method. JD, ZD, WM and DL participated in the
465 discussions. WL, WT and WF provided resources. All the authors have made substantial contributions to this
466 article.

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