

1 An investigation on hygroscopic properties of 15 black carbon
2 (BC)-containing particles from different carbon sources: Roles
3 of organic and inorganic components
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Abstract

The hygroscopic behavior of black carbon (BC)-containing particles (BCPs) has a significant impact on global and regional climate change. However, the mechanism and factors controlling the hygroscopicity of BCPs from different carbon sources are not well understood. Here, we systematically measured the equilibrium and kinetics of water uptake by 15 different BCPs (10 herb-derived BCPs, 2 wood-derived BCPs, and 3 soot-type BCPs) using gravimetric water vapor sorption method combined with in-situ diffuse reflectance infrared Fourier transform spectroscopy (DRIFTS). In the gravimetric analysis, the sorption/desorption equilibrium isotherms were measured under continuous-stepwise water vapor pressure conditions, while the kinetics was measured at a variety of humidity levels obtained by different saturated aqueous salt solutions. The equilibrium water uptake of the tested group of BCPs at high relative humidity ($> 80\%$) positively correlated to the dissolved mineral content ($0.01\text{--}13.0\text{ wt\%}$) ($R^2 = 0.86$, $P = 0.0001$), the content of the thermogravimetrically analyzed organic carbon (OC_{TGA} , $4.48\text{--}15.25\text{ wt\%}$) ($R^2 = 0.52$, $P = 0.002$), and the content of the alkali-extracted organic carbon (OC_{AE} , $0.14\text{--}8.39\text{ wt\%}$) ($R^2 = 0.80$, $P = 0.0001$). In contrast, no positive correlation was obtained with the content of total organic carbon or elemental carbon. Among the major soluble ionic constituents, chloride and ammonium were each correlated with the equilibrium water uptake at high relative humidity. Compared with the herbal BCPs and soot, the woody BCPs had much lower equilibrium water uptake, especially at high relative humidity, likely due to the very low dissolved mineral content and OC content. The DRIFTS analysis provided generally consistent results at low relative humidity. The kinetics of water uptake (measured by pseudo-second-order rate constant) correlated to the content of OC_{TGA} and OC_{AE} as well as the content of chloride and ammonium at low relative humidity

(33%), but to the porosity of BCPs at high relative humidity (94%). This was the first study to show that BCPs of different types and sources had greatly varying hygroscopic properties.

1. Introduction

Black carbon (BC) refers to a collective term of refractory and insoluble aggregates of small carbon spherules generated from incomplete combustion of biomass and fossil fuels (Bond et al., 2013). BC-containing particles (BCPs) are ubiquitous in the atmosphere and are a major component of atmospheric carbonaceous aerosols (Schwarz et al., 2008). Due to the strong ability to absorb visible light (Yuan et al., 2015), BCPs cause positive radiative forcing effects on climate and are considered an important factor driving global warming (Matthews et al., 2009). Once immersed into cloud droplets, BCPs can also facilitate water evaporation and cloud dispersion via enhanced absorption of solar radiation, and thus produce indirect radiative forcing effects (Powelson et al., 2014). Additionally, the large specific surface area of BCPs creates a potential for heterogeneous reactions with trace gases (such as volatile halocarbons) in the atmosphere (Qiu et al., 2012), therefore heavily impacting atmospheric chemistry and air quality. Hygroscopicity is a key determinant of physical, chemical, and optical properties of BCPs by changing particle size, phase state, and quality and morphological development, which in turn affect aerosol radiation effect, formation of cloud and ice nuclei, and heterogeneous chemical reactions (Bond et al., 2013; Liu et al., 2018). Furthermore, the hygroscopicity of BCPs is an important factor contributing to the risk of human respiratory infections, cardiovascular diseases, and other infectious diseases (Haddrell et al., 2015).

Inherited from parent carbon source or produced from combustion process, large content of inorganic impurities are often present in BCPs, consisting of a variety of

amorphous or crystalline salts (sulfates, chlorides, etc.) as well as semi-crystalline minerals (such as silica) (Stanislav et al., 2013). Despite the relatively low content, the inorganic components play in BCPs a significant role in water uptake, depending on their types, contents, and mixing ratios (Lewis et al., 2009). As the factors and processes governing the hygroscopic deliquescence of inorganic salts are very complicated (Reid et al., 2005; Zhang et al., 2012), it is of a great challenge to assess the contribution of a specific salt to the overall hygroscopicity of BCPs, and thus, its role is still controversial. Previous studies suggested that KCl was responsible for the high hygroscopicity of BCPs produced by fresh biomass burning (Posfai et al., 2003), while the presence of K_2SO_4 or KNO_3 caused the low hygroscopicity of BCPs produced by aged biomass burning (Li et al., 2003).

The organic components in BCPs consist of graphitized elemental carbon (EC) and non-condensed, amorphous organic carbon (OC) (Lian and Xing, 2017). The contribution of EC to the overall hygroscopicity of BCPs is considered low due to the very high hydrophobicity (Seisel et al., 2005). The role of OC in the hygroscopic growth of BCPs is intricate and debatable. The positive effect of OC is mainly attributed to water absorption by the oxygen-containing functional groups (Fletcher et al., 2007; Suda et al., 2014). The negative effect of OC is suggested to stem from the impeded mass transfer process of water molecules by formation of coatings on hygroscopic minerals or inhomogeneous morphology inside the particle (Sjogren et al., 2007; Stemmler et al., 2008). In addition to the total content, the molecular weight, water solubility, surface tension, and type and content of functional groups of OC were all found to influence the overall hygroscopicity of BCPs. Moreover, the effect of OC on BCPs hygroscopicity is further complicated by the formation of organic minerals (presumably through strong covalent bonds) (Archanjo et al., 2014; Reid et

al., 2005; Zuend et al., 2011).

The carbon sources for BCPs released into the atmosphere are expected to be highly diversified and cover a wide range of plant biomass, coals, and refined oil products, although their quota can hardly be accurately assessed (Andreae and Gelencser, 2006). The chemical compositional, structural, and morphological properties of BCPs depend significantly on carbon sources and combustion conditions (Xiao et al., 2018). For instance, crop residue-derived biochar often has higher mineral content than wood-derived charcoal, while biochar formed at higher pyrolysis temperatures generally have higher aromaticity, specific surface area, and pore volume but lower polarity than biochar formed at lower temperatures (Wei et al., 2019). Compared with char and charcoal, soot (another type of BCPs) produced from fossil fuel combustion is comprised of more regular shaped, chain-like agglomerates of primary particles, which consist of perturbed graphitic layers oriented concentrically in an onion-like fashion (Nienow and Roberts, 2006). Previous studies on hygroscopicity of BCPs have mainly focused on wood-derived BCPs (Carrico et al., 2010; Day et al., 2006), whereas BCPs from other carbon sources has been largely overlooked. It remains unclear whether BCPs from different carbon sources would differ significantly in hygroscopicity.

Herein we systematically investigated the equilibrium and kinetics of water uptake by 15 different BCPs derived from wood, herb, coal, and diesel at varying relative humidity (RH) levels by gravimetric sorption and in-situ diffuse reflectance infrared Fourier transform spectroscopy (DRIFTS). The chemical, compositional, and structural properties of the tested group of BCPs were thoroughly characterized to unveil the key factors controlling the hygroscopic properties.

2. Experimental methods

2.1. Preparation of BCPs

A total of 15 BCPs were tested, including 10 herb-derived BCPs, 2 wood-derived BCPs, and 3 soot-type BCPs. The herbs used for preparation of herbal BCPs were amaranth, peanuts, pea, grass, rice, wheat, corn, millet, sorghum (Nantong, Jiangsu Province, China), and bamboo (Lishui, Zhejiang Province, China), and the woods used for preparation of woody BCPs were red pine and poplar (Lishui, Zhejiang Province, China). The dried and dehydrated biomass was pulverized into a fine powder using a high-speed pulverizer (FW 100, Tianjin Taisite Instrument, China), and pyrolyzed in a muffle furnace under an oxygen-limited conditions. The oven temperature was programmed to increase from 20 to 400 °C in 2 h and maintained at 400 °C for 3 h. The Weifu diesel soot was purchased from Wuxi Weifu Automotive Diesel System Co., Ltd. (Jiangsu Province, China). According to the information provided by the manufacturer, the soot was produced under laboratory conditions by burning diesel (Diesel # 5, China) at 1000 °C and was collected by a diesel particulate filter (NGK-6000YE) from the exhaust stream at a carbon deposition temperature of 250 °C. The diesel engine soot was collected in outdoor conditions from the freshly discharged exhaust particles on the tailpipe of a diesel truck (Dongfeng, CY4100, 2015, China) equipped with a diesel engine (CY4100Q, 3.7 L, Diesel # 0) and a diesel particulate filter (BST-5L-QCD). The household soot was collected on the inner wall of a stove chamber (Linkou County, Heilongjiang Province, China) 1 hour later after burning of coal and wood for winter cooking and heating under limited oxygen exposure. The obtained samples of BCPs were further ground to pass a 100-mesh sieve (0.15 mm) and stored sealed in a brown glass bottle at 4 °C.

2.2. Characterization of BCPs

Elemental analysis (EA) was performed using a Vario micro cube elemental analyzer

(Elementar, Hanau, Germany). Surface elemental compositions were measured by X-ray photoelectron spectroscopy (XPS) (PHI 5000 VersaProbe, UIVAC-PHI, Japan). Mineral compositions were measured by X-ray fluorescence (XRF) (ARL-9800, ARL Corporation, Switzerland). Fourier-transform infrared (FTIR) spectra were recorded on a Bruker Tensor 27 Karlsruhe spectrometer (Germany) using KBr pellets in the range of 400 to 4000 cm^{-1} . X-ray diffraction (XRD) spectra were recorded on an AXS D8 Advance spectrometer (Germany) using Cu $K\alpha$ radiation at a 2θ angle ranging from 5 to 70°. Raman spectra were collected on a Horiba Jobin Yvon LabRam HR-800 spectrometer equipped with a 514 nm laser (France). N_2 adsorption isotherms to the 15 BCPs were obtained on a Micrometrics ASAP 2020 (Micromeritics Instrument Co., Norcross, GA, USA) apparatus at -196 °C (77 K).

Three different methods, thermogravimetric analysis (TGA), alkali extraction, and water extraction, were explored to quantify the content of OC in BCPs (referred to as OC_{TGA} , OC_{AE} , and OC_{WE} , respectively). The content of OC_{TGA} was measured as the weight loss during the heating of BCPs from 30 to 300 °C at a ramp of 10 °C per minute in a nitrogen flow (Han et al., 2013) using a TGA 8000 analyzer (PerkinElmer, USA). To measure the content of OC_{AE} , the sample of BCPs was mixed with 0.1 M NaOH at a solid-to-solution ratio of 1: 10 (w/w) and magnetically stirred for 12 h, followed by filtration through a 0.45- μm filter membrane (Pall, USA) (Song et al., 2002). The filtrate was collected and acidified (pH 1.0 with 6 M HCl), and the total organic carbon (TOC) content of the sample was measured by a TOC analyzer (TOC-5000A, Shimadzu, Japan). For the three selected BCPs (Grass, Wheat, and Household soot), the precipitates formed by acidification were further separated by centrifugation and dialyzed in deionized (DI) water by dialysis bag (500 Da, Union Carbide, USA) until no chloride ion was detected by AgNO_3 , and then freeze-dried.

The elemental compositions of the three prepared OC_{AE} were measured by EA. To measure the content of OC_{WE} and dissolved minerals, the suspended BCPs in DI water (BCPs-to-water ratio of 1: 10, w/w) was sonicated in a water bath for several minutes, and the mixture was filtered through a 0.45-μm membrane. The suspension/sonication procedure was repeated for 6 times. The filtrate was collected and subjected to TOC analysis to obtain the content of OC_{WE}. The concentrations of ionic constituents (Cl⁻, NO₃⁻, PO₄³⁻, SO₄²⁻, F⁻, COO⁻, C₂O₄²⁻, Na⁺, NH₄⁺, K⁺, Mg²⁺, Ca²⁺, Al³⁺) in the filtrate were measured using a Dionex ICS-1100 ion chromatography (Thermo Scientific, USA). The cations were eluted using 20 mM methanesulfonic acid on a Dionex IonPac CS12A column (4 × 250 mm), while the anions were eluted using an eluent of 4.5 mM Na₂CO₃ and 0.81 mM NaHCO₃ on a Dionex IonPac AS14A column (4 × 250 mm). The filtrate was further freeze-dried and baked at 600 °C for 6 h to remove organic components. The remaining ash was weighed to determine the content of dissolved minerals in BCPs. A portion of the ash was extracted three times using DI water at a solid-to-solution ratio of 1:10 (w/w) under sonication, and the salinity of the extract was measured by a ST3100C conductivity meter (OHAUS, USA). All reagents and chemicals used were of analytical reagent grade.

2.3. Measurement of hygroscopicity of BCPs

The hygroscopicity of BCPs at varying RH was measured by gravimetric method combined with in-situ DRIFTS. The water vapor sorption/desorption isotherms to BCPs under a range of continuous-stepwise water vapor pressures were acquired on a 3H-2000 PW Multi-stations Gravimetric Method Steam Adsorption Instrument (Beijing, China) at 25 °C using an approach similar to that in previous studies (Gu et al., 2017). The instrument consists of two main parts: a balance chamber to determine

the sample mass to an accuracy $\pm 1 \mu\text{g}$ and a humidity chamber to regulate the water vapor pressure to the desired value as monitored online by a pressure sensor. Prior to testing, the BCPs (about 10 mg) were dried at 70°C under vacuum for 12 h to remove residual water vapor and set a baseline for comparison of water vapor adsorption behaviors among different samples. The amount of water sorbed to BCPs was monitored as the mass difference before and after sorption. The amount of water sorbed to the sample tube was measured and showed to be negligible ($< 0.05\%$ of the amount of water sorbed to BCPs). The water vapor pressures ranging from 10 to 94% RH were applied to the sorption isotherm branch in a stepwise increasing sequence and to the desorption isotherm branch in a stepwise decreasing sequence.

The kinetics of water sorption to BCPs was measured on a 100 mm closed quartz chamber (Jiangsu Province, China) using a gravimetric method similar to that in previous studies (Yuan et al., 2014). Approximately 100 mg of BCPs was dried at 70°C under vacuum for 12 h, weighed in a 10-mL beaker, and placed in a chamber under controlled humidity conditions based on different saturated aqueous salt solutions according to ASTM E104-02 (2007). The saturated solutions of CH_3COOK , MgCl_2 , K_2CO_3 , LiNO_3 , NaCl , KCl , and KNO_3 provided RH of 23%, 33%, 43%, 47%, 75%, 84%, 94%, respectively at 25°C . The sample was continuously weighed and recorded over a period of time (48 h for low humidity and 96 h for high humidity) to monitor the amount of sorbed water. The RH was monitored in real time using a Honeywell HIH4000 hygrometer (USA) with measurement variance was less than 5%. Sorption equilibrium was reached in the late stage of the experiment as evidenced by the stabilized constant value of sample mass. In addition to kinetic data, sorption isotherms were also collected for the seven selected RH levels using the measured mass under equilibrium conditions.

Samples of BCPs equilibrated at different RH levels were characterized by in situ DRIFTS using a Bruker Tensor 27 spectrometer equipped with a high-sensitivity mercury-cadmium-telluride (MCT) detector working under liquid N₂ conditions and a chamber fitted with ZnSe windows (Harrick Scientific, USA). About 10 mg of BCPs pre-dried at 70 °C under vacuum for 12 h was transferred to the chamber which was connected to a gas feeding system. The chamber was sealed and purged with high-purity N₂ at a flow rate of 100 mL per minute for at least 3 h to remove pre-adsorbed gases on BCPs and to minimize the interference of environmental CO₂. The humidity in the chamber was regulated by mixing high-purity N₂ and saturated water vapor at 25 °C with varying ratios and monitored in real time by a hygrometer (Vaisala Humitter, Australia). The sample was equilibrated with the gas mixture in the chamber for at least 30 minutes to reach sorption equilibrium based on pre-determined kinetics. The spectra were acquired by co-adding and averaging a plurality of 500 scans with a resolution of 4 cm⁻¹ (Song and Boily, 2013). The amount of water sorbed to BCPs was monitored by the integrated intensity of the O-H stretching region from 2750 to 3660 cm⁻¹ (Ghorai et al., 2011).

3. Results and discussion

3.1. Characteristics of BCPs

Bulk elemental compositions by EA and surface elemental compositions by XPS are summarized in Table S1. The bulk elemental compositions of all BCPs were dominated by C and O, together accounting for 54%–96% of the total. However, the bulk C, O compositions differed significantly among the 15 BCPs, ranging from 32 to 76% for C and from 16 to 69% for O. With the exception of the woody BCPs, the differences were apparent within each category of the herbal BCPs and the soot. Similarly, the surfaces of all the 15 BCPs were dominated by C and O; however, the

240 differences of the surface C, O compositions among the 15 BCPs were much smaller
241 compared to the bulk C, O compositions. Besides C and O, EA detected low amounts
242 of N ($< 3.7\%$) and S ($< 1.8\%$), and XPS detected low amounts of N ($< 4.3\%$), Si ($<$
243 5.6%), and S ($< 0.6\%$). The contents of oxygen-containing groups in the 15 BCPs
244 were qualitatively compared by the FTIR spectra (Figure S1). All the tested BCPs
245 except Weifu diesel soot showed characteristic peaks of esters (1700 cm^{-1}), ketones
246 ($1613, 1100\text{ cm}^{-1}$), and phenols (1270 cm^{-1}) (Keiluweit et al., 2010), generally with
247 larger peak intensities observed for herbal BCPs and household soot.

248 Table 1. Chemical, compositional, and pore properties of different BCPs.

BCPs	OC			EC ^d (wt%)	Dissolved minerals (wt%)	Total Porosity ^e (m ³ g ⁻¹)	SSA ^f (m ² g ⁻¹)
	OC _{TGA} ^a (wt%)	OC _{AE} ^b (wt%)	OC _{WE} ^c (wt%)				
Amaranth	6.24	2.6	1.75	25.84	10.8	0.004	0.314
Grass	7.24	2.37	1.01	51.56	4.8	0.008	5.587
Peanuts	7.45	1.78	0.8	41.86	4.2	0.002	0.192
Pea	9.59	1.98	0.09	54.48	3.6	0.005	4.679
Rice	6.81	0.6	0.11	48.06	0.6	0.023	31.88
Wheat	8.25	1.82	0.37	42.65	5.8	0.01	7.382
Millet	9.41	1.97	0.93	32.25	8	0.023	8.319
Corn	6.55	0.14	0.32	46.47	1.8	0.028	28.6
Sorghum	9.09	1.12	0.62	55.26	4.8	0.001	0.192
Bamboo	6.84	0.23	0.12	61.7	0.6	0.029	51.94
Red pine	7.4	0.2	0.05	62.59	0.01	0.032	64.24
Poplar	7.58	0.19	0.09	64.22	0.6	0.071	107.6
Diesel engine soot	9.57	1.4	0.78	27.37	3.6	0.021	6.119
Weifu diesel soot	4.48	0.57	0.13	71.98	3.4	0.484	194.6
Household soot	15.25	8.39	2.24	21.83	13	0.012	7.79

249 ^aContent of organic carbon determined by TGA. ^bContent of alkali-extracted organic
250 carbon determined by TOC analysis. ^cContent of water-extracted organic
251 carbon determined by TOC analysis. ^dDetermined by subtracting OC_{TGA} content from
252 total organic carbon content by EA. ^eTotal pore volume determined by N₂ adsorption
253 at 0.97 atmosphere pressure. ^fSpecific surface area determined by the BET method.

254 Table 1 summarizes the contents of OC (OC_{TGA}, OC_{AE}, and OC_{WE}) of the 15 BCPs
255 by TGA, alkali extraction, and water extraction, respectively. For a given BCPs, the
256 contents of the three types of OC differed pronouncedly, with an increasing order of
257 OC_{WE} < OC_{AE} < OC_{TGA}. The OC content also differed within the tested group of
258 BCPs, ranging from 0.05 to 2.24wt% for OC_{WE}, from 0.14 to 8.39wt% for OC_{AE}, and
259 from 4.48 to 15.25wt% for OC_{TGA}. Compared with the EC (graphitized carbon), the
260 three types of OC are non-condensed, amorphous, and more rich in oxygen-containing
261 functional groups, which was evidenced by the fact that the OC_{AE} from the three
262 selected BCPs (Grass, Wheat, and Household soot) had markedly higher bulk

compositions of O (results presented in Table S2). The EC content in BCPs was roughly assessed by subtracting the OC_{TGA} content from the total organic carbon content measured by EA (results presented in Table 1). The calculated EC content negatively correlated with the OC_{AE} content ($R^2 = 0.43$, $P = 0.0079$) for the examined group of BCPs. This relationship was reasonable as EC was comprised of mature, thermodynamically stable graphitized carbons, while OC was comprised of the less mature and less aromatic constituents remaining after pyrolysis. Except for Weifu diesel soot, the two woody BCPs had the highest EC, but the lowest OC_{AE} and OC_{WE} among the 15 BCPs.

The relative abundance of EC in BCPs was also assessed by Raman spectroscopy (Figure S2). The spectra of all the tested BCPs were dominated by a D band at 1340 cm^{-1} and a G band at 1580 cm^{-1} , which were ascribed to carbon network defects and the E_{2g} mode of the graphitized carbon, respectively (Pimenta et al., 2007). Thus, the ratio of these two bands (I_D/I_G) was inversely proportional to the in-plane crystallite size of graphitized carbons of BCPs (Cancado et al., 2006). The I_D/I_G ratio of the woody BCPs (0.51–0.59) was less than those of the herbal BCPs (0.88–1.09) and the soot (0.77–1.12) (Table S3), suggesting larger sizes of graphitized carbons in the woody BCPs, which was consistent with the results of OC compositions.

The contents of dissolved minerals of the 15 BCPs are listed in Table 1, and their salinities in water extracts are listed in Table S5. The two woody BCPs had the lowest contents of dissolved minerals and salinities, while these contents in herbal BCPs and soot were higher and varied greatly. The mineral compositions characterized by XRF are listed in Table S6. Si-, K-rich minerals were the two major inorganic constituents in the herbal BCPs and woody BCPs. Moreover, these two types of BCPs generally contained trace amounts of S-, Cl-, Ca-, P-, Mg-, Na-, Fe-, and Al-minerals, with

lower contents observed for the woody BCPs. The three soot had very different mineral compositions. Household soot was dominated by S-, Ca-, Si-, and Cl-minerals, Diesel engine soot was dominated by S-, Ca-, and Fe-minerals, while Weifu diesel soot contained negligible mineral compositions. As reflected by the observed characteristic peaks and associated peak intensities in the XRD spectra (Figure S3), the herbal BCPs and Household soot contained more mineral species with higher contents than other BCPs, whereas the two woody BCPs and Weifu diesel soot contained the least species and contents of minerals. Potassium salts, amorphous silica, and sulfates were the major minerals in the herbal BCPs. Soot had the largest content of sulfates among the tested BCPs. According to the ion chromatograph analysis (results presented in Figure S4 and Table S7), the major water-extracted cationic species from the tested BCPs were NH_4^+ , K^+ , and Ca^{2+} , and the major anionic species were SO_4^{2-} , Cl^- , and $\text{C}_2\text{O}_4^{2-}$. The herbal BCPs had high contents of K^+ , $\text{C}_2\text{O}_4^{2-}$, and Cl^- , while the soot had high content of SO_4^{2-} .

The Brunauer–Emmett–Teller (BET) specific surface area and total porosity measured by N_2 adsorption are also summarized in Table 1. A huge disparity of specific surface area was shown among the 15 BCPs and among the BCPs within each category, ranging from 6 to $200 \text{ m}^2 \text{ g}^{-1}$ for the soot, from 60 to $110 \text{ m}^2 \text{ g}^{-1}$ for the woody BCPs, and from 0.1 to $52 \text{ m}^2 \text{ g}^{-1}$ for the herbal BCPs. The micro- and mesoporosity of the 15 tested BCPs are summarized in Table S4. The herbal BCPs and woody BCPs were dominated by micropores (pore size $< 2 \text{ nm}$), which accounted for more than 50% of the total pore volume. Alternatively, mesopores ($50 \text{ nm} > \text{pore size} > 2 \text{ nm}$) were the main pore structure of the soot, accounting for more than 61% of the total pore volume.

3.2. Hygroscopic properties of BCPs

Equilibrium water uptake. Figure 1 displays sorption and desorption isotherms of water vapor with BCPs plotted as equilibrium water uptake (mg g^{-1}) by unit mass of BCPs under continuous-stepwise water vapor pressure conditions. Figure S5 displays the equilibrium sorption isotherms at selected humidity levels obtained by using saturated aqueous salt solutions. Under similar humidity conditions (80% and 84%), the water uptake by the 15 BCPs was very close between these two humidity regulation methods (Figure S6), reflecting their technical validity. The woody BCPs showed very different sorption isotherm patterns from the herbal BCPs and soot. First, the water sorbing ability of the woody BCPs was much lower. The maximum water uptake observed at the highest RH (94%) was approximately 65 mg g^{-1} by the woody BCPs, but was more than 400 mg g^{-1} for the strongest sorbing herbal BCPs and soot. Second, much larger water-uptake disparities were observed within the group of herbal BCPs and the soot group than within the group of woody BCPs. Additionally, over the examined RH range (10–94%), the water uptake by the woody BCPs increased slowly and linearly with the RH; however, for the herbal BCPs and soot, the water uptake increased more rapidly with the RH, especially under high humidity conditions ($\text{RH} > 70\%$).

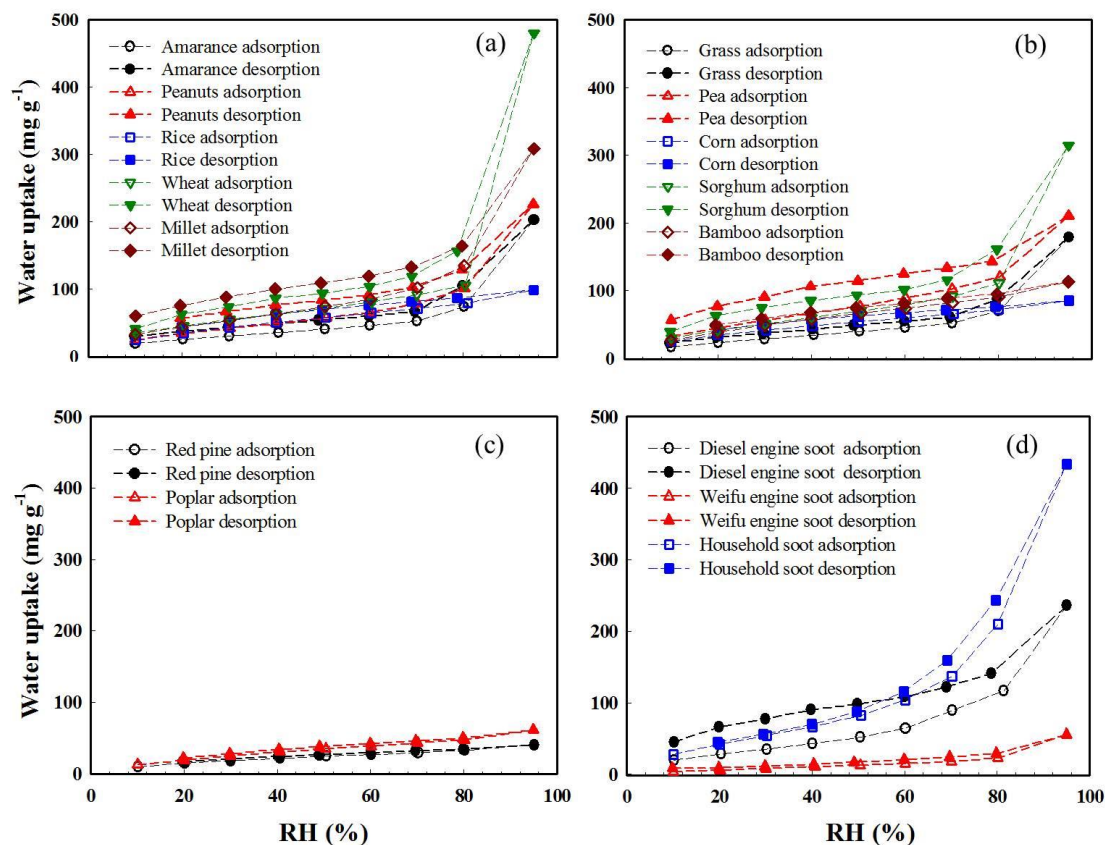


Figure 1. Sorption and desorption isotherms of water vapor plotted as water uptake (mg g^{-1}) vs. relative humidity (RH, %) at equilibrium for different BCPs. (a) Subgroup 1 of herbal BCPs. (b) Subgroup 2 of herbal BCPs. (c) Woody BCPs. (d) Soot. Data were collected with single-point.

To better understand the underlying mechanisms and factors controlling the hygroscopic properties of BCPs, linear relationships were built between the equilibrium water uptake and a wide variety of compositional and pore property parameters of the whole group of BCPs. Figure 2 displays regression relationships with contents of OC_{TGA} , OC_{AE} , EC, dissolved minerals, major ionic species (NH_4^+ , Cl^- , SO_4^{2-} , and $\text{C}_2\text{O}_4^{2-}$), and total porosity, respectively at 94% RH. The regression relationships at 23% RH were presented in Figure S7. The regression accuracy (R^2 and P) values at the 7 different RH levels (23, 33, 43, 47, 75, 84, 94%) are summarized in Table S8. Good positive correlations existed between the water uptake

and the contents of OC_{TGA}, OC_{AE}, and dissolved minerals under high humidity conditions (Figure 2a-d). The highest regression accuracy values obtained were $R^2 = 0.82$, $P < 0.0001$ for OC_{TGA} at 84% RH, $R^2 = 0.80$, $P = 0.001$ for OC_{AE} at 94% RH, and $R^2 = 0.86$, $P = 0.0001$ for dissolved minerals at 94% RH. However, the correlations with these constituents became much weaker under low humidity conditions ($R^2 = 0.10$ – 0.32 , $P = 0.247$ – 0.028 at RH = 23%). It can be concluded that the hygroscopicity of herbal BCPs and soot under high humidity conditions was mainly controlled by the contents of OC and dissolved minerals. On the other hand, the low water sorbing ability of the woody BCPs was due to the very low contents of these constituents. As indicated by the elemental analysis results (Table S2), the OC constituents in the three representative BCPs (Grass, Red pine, and Household soot) contained large amounts of oxygen-containing groups, which expectedly had very high hygroscopicity (Xiao et al., 2013). The very strong water retention ability of dissolved minerals in BCPs was understandable due to the strong hydration of mineral surfaces and ionic constituents. No correlation was observed between the water uptake of BCPs and the total organic carbon content within the whole examined RH range. Notably, a negative correlation was observed with the EC content (Figure 2c), especially under high humidity conditions ($R^2 = 0.54$, $P = 0.0019$ at 94% RH). Compared with the OC in BCPs, the EC was comprised mainly of hydrophobic fused aromatic hydrocarbons and had much lower amounts of oxygen-containing groups, resulting in the very low water sorbing ability.

The above one-factor linear correlation analysis indicated that at high RH the hygroscopicity of BCPs was dominated by OC and dissolved minerals. Consistently, a better positive correlation ($R^2 = 0.90$, $P < 0.0001$) could be obtained between the equilibrium water uptake and the contents of OC_{TGA} and dissolved mineral by binary

linear regression for the tested group of BCPs (Figure S8). Interestingly, the correlation relationship became poorer as the RH level gradually decreased, and the worst correlation was shown at 23% RH ($R^2 = 0.21$, $P = 0.028$). The results suggested that the hygroscopicity of BCPs at low RH was controlled by different factors.

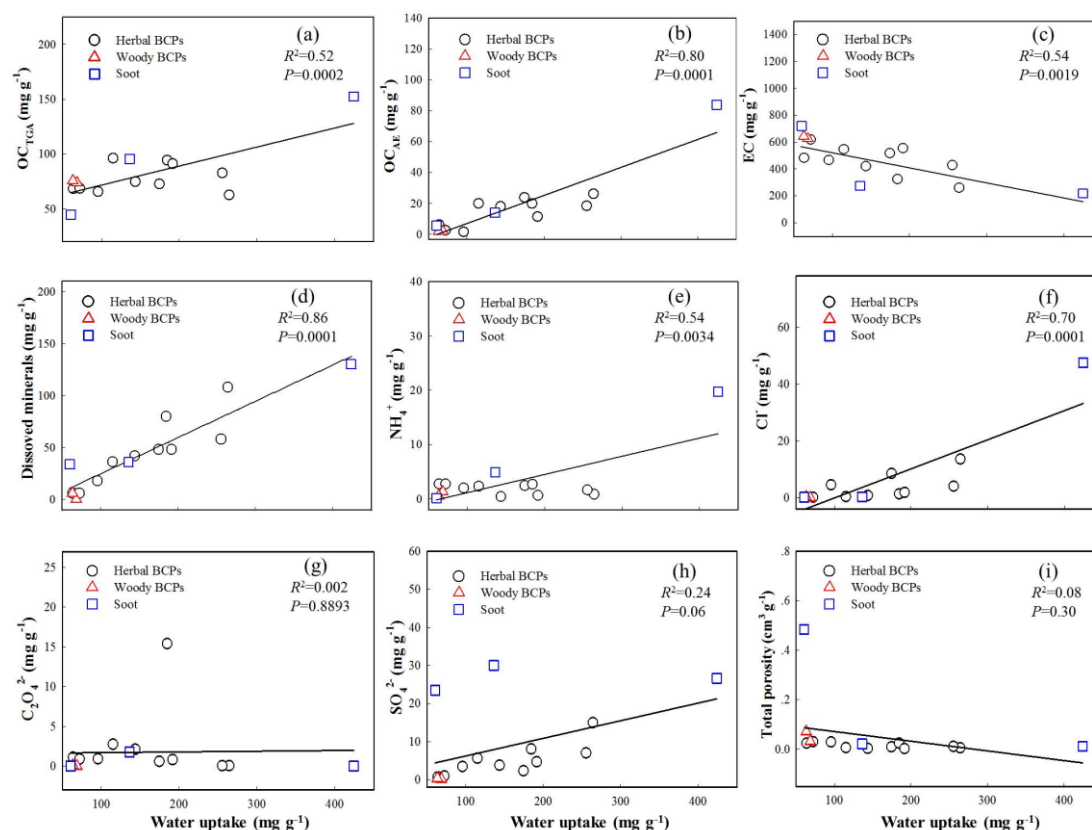


Figure 2. Relationships between equilibrium water uptake (mg g⁻¹) vs. compositional and pore property parameters for the group of BCPs at 94% relative humidity. (a) TGA-measured organic carbon (OC_{TGA}). (b) Alkali-extracted organic carbon (OC_{AE}). (c) Elemental carbon (EC). (d) Dissolved minerals. (e) Ammonium (NH₄⁺). (f) Chloride (Cl⁻). (g) Oxalate (C₂O₄²⁻). (h) Sulfate (SO₄²⁻). (i) Total porosity.

At 94% RH, relatively good positive correlations were observed with NH₄⁺ and Cl⁻ ($R^2 = 0.50$ – 0.70 , $P = 0.0001$ – 0.0034), but not with SO₄²⁻ ($R^2 = 0.24$, $P = 0.06$) or C₂O₄²⁻ ($R^2 = 0.002$, $P = 0.89$) (Figure 2e-h). No correlation ($R^2 = 0.08$, $P = 0.30$) was observed with the total porosity of BCPs at 94% RH (Figure 2i). Consistently,

previous studies reported that chloride salts in biomass burning aerosols had high hygroscopicity (Jing et al., 2017; Posfai et al., 2003). The poor correlation observed for SO_4^{2-} was ascribed to the low hygroscopicity of CaSO_4 and K_2SO_4 , as evidenced by their very high deliquescent relative humidity (96–97%) (Freney et al., 2007; Preturlan et al., 2019). It is noteworthy that the content of SO_4^{2-} positively correlated with the contents of Ca^{2+} ($R^2 = 0.74$, $P < 0.0001$) and K^+ ($R^2 = 0.69$, $P = 0.0008$) for the tested BCPs. On the other hand, the poor correlation observed for $\text{C}_2\text{O}_4^{2-}$ was likely due to the formation of less water-soluble compounds (e.g., CaC_2O_4 and $\text{H}_2\text{C}_2\text{O}_4$) that might depress the hygroscopicity (Buchholz and Mentel, 2008).

The positive correlations observed with NH_4^+ and Cl^- at 94% RH disappeared at 23% RH. Alternatively, a relative good negative correlation ($R^2 = 0.42$, $P = 0.0095$) with the total porosity was shown at 23% RH (Figure S7i). Similarly, a weak negative correlation ($R^2 = 0.21$, $P = 0.083$) was shown with the EC content at 23% RH. The consistency between the total porosity and the EC content was likely due to the dominant role of graphitized carbons (EC) in forming rigid pore structures of BCPs (Han et al., 2014). On the contrary, a weak positive correlation ($R^2 = 0.32$, $P = 0.028$) was observed between the water uptake and the OC_{TGA} content at 23% RH, indicating that the OC constituents were an important factor contributing to the overall water uptake by BCPs under low humidity conditions.

As can be seen from the sorption-desorption isotherms in Figure 1, some herbal BCPs (Amarance, Peanuts, Wheat, Millet, Pea, and Sorghum) and soot (Diesel engine soot and Household soot) showed strong hysteresis effect (irreversible sorption), whereas none of the woody BCPs showed hysteresis effect. Irreversible sorption would lower the release of sorbed water molecules from BCPs in the atmosphere when the RH changes from a high level to a low level. The observed hysteresis effect

of herbal BCPs and soot likely stemmed from their relatively high contents of OC and/or dissolved minerals (such as Wheat and Household soot). Sorbing water molecules could cause strong and irreversible hydration of organic acids (Petters et al., 2017) and dissolution or phase change of minerals (Adapa et al., 2018), consequently leading to hysteresis effect due to non-identical structures of BCPs between the sorption and desorption branches even at the same RH. The negligible hysteresis effect observed on the two woody BCPs could be attributed to their very low contents of OC and dissolved minerals.

The equilibrium water uptake by BCPs was further investigated by DRIFTS. The spectra of representative BCPs (Grass, Red pine, and Household soot) at varying RH are presented in Figure 3a-c. Figure 3d compares the water uptake at 23% RH monitored by the integrated intensity of the O-H stretching region from 2750 to 3660 cm^{-1} (Ghorai et al., 2011), along with the water uptake measured by the multi-station gravimetric method for 8 selected BCPs. The identified bands of sorbed water molecules included a combination mode of symmetric stretch around 3423 cm^{-1} and asymmetric stretch around 3253 cm^{-1} (Gustafsson et al., 2005). The broad feature peak centered at 2100 cm^{-1} was assigned to a combined band of bending, libration, and hindered translation modes of water, while the peak centered at 1640 cm^{-1} was attributed to the bending mode of water (Ma et al., 2010). The intensities of these peaks/bands increased with increasing RH. As assessed by the integrated intensity of the O-H stretching region (see insets in Figure 3a-c), the water uptake by Grass and Household soot increased gradually with RH from 12 to 80%; however, the water uptake by Red pine rapidly reached saturation at about 28% RH, and kept constant when the RH was further increased. With the exception of Weifu diesel soot, the disparity pattern of water uptake by the 8 selected BCPs at 23% RH monitored by

DRIFTS was similar to that monitored by the multi-station gravimetric method (Figure 3d). However, the disparities were very large between these two methods at 80% RH (Figure S9), which was probably because the saturated effect encountered in detection of sorbed water molecules by FTIR (Gustafsson et al., 2005) became worse under high humidity conditions. Moreover, the DRIFTS signals of sorbed water molecules might be influenced by the distribution of sorption sites (e.g., minerals vs. OC and exterior vs. interior), and the caused effects might lead to larger deviations under high humidity conditions.

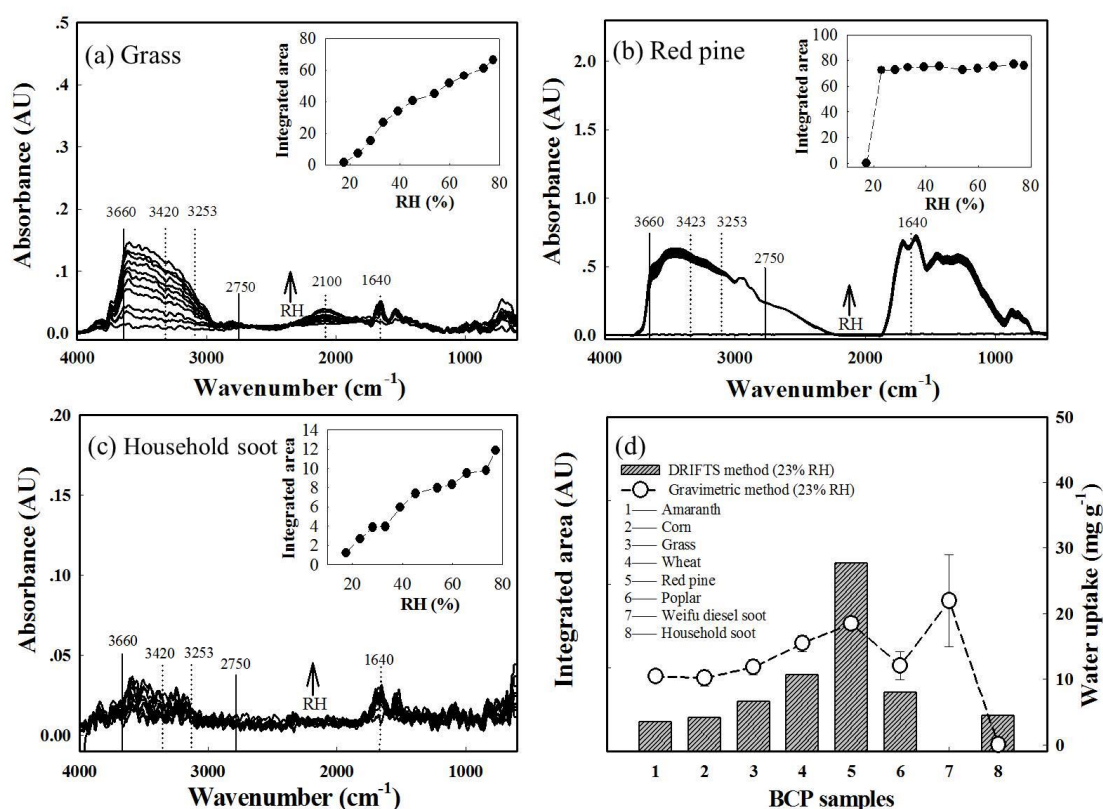


Figure 3. Diffuse reflectance infrared Fourier transform spectroscopy (DRIFTS) characterization of equilibrium water uptake by BCPs. (a-c) Spectra for the three representative BCPs (Grass, Red pine, and Household soot) equilibrated at varying relative humidity (RH) levels. (d) Comparison of equilibrium water uptake measured as integrated area of O-H stretching region ($2750\text{--}3660\text{ cm}^{-1}$) between 8 selected

BCPs at 23% relative humidity. Error bars represent standard variations from duplicate samples. Insets in subfigures (a-c) present water uptake measured as integrated area of O-H stretching region against RH.

Kinetic water uptake. Figure 4 displays the water vapor sorption kinetics to the 15 BCPs at 94% RH obtained by saturated aqueous salt solutions. The sorption kinetics at 33% RH was presented in Figure S10. The two woody BCPs exhibited similar kinetics curves; however, the herbal BCPs and soot showed very different kinetic patterns within each group. The kinetic data were fitted to the pseudo-first-order and pseudo-second-order models, $dq_t/dt = k_1 (q_e - q_t)$ and $dq_t/dt = k_2 (q_e - q_t)^2$, respectively, where q_t was the sorbed concentration at time t , q_e was the equilibrium sorbed concentration, and k_1 and k_2 were the pseudo-first-order and pseudo-second-order rate constants, respectively. Note that these two kinetic models were applied only for quantitative comparison of apparent sorption kinetics among different BCPs, but not for illustration of sorption mechanisms. The fitting parameters (q_e , k_1 , k_2) for the three selected RH levels (33, 47, and 94%) are summarized in Tables S9-S10. The pseudo-second-order model ($R^2 > 0.97$) fits the data better than the pseudo-first-order model ($R^2 = 0.80\text{--}0.99$). The calculated k_2 differed greatly among the BCPs within the group of herbal BCPs and the soot group, but was very close between the two woody BCPs. For a given BCPs, the k_2 at a lower RH level was significantly larger than that at a higher RH level. Similar results were reported in previous studies on sorption kinetics of water vapor to activated carbon (Ohba and Kaneko, 2011; Ribeiro et al., 2008). Under low humidity conditions, sorption of water vapor mainly occurs at the active, high-energy binding sites, and the sorption kinetics is fast; alternatively, under high humidity conditions, sorption is governed by the slow pore-filling/condensation process of water molecules within the pores of activated carbon via formation of

water clusters around the water molecules already sorbed at the active sites (Nguyen and Bhatia, 2011; Rosas et al., 2008). Due to the small molecular size (0.0958×0.151 nm, ChemDraw 3D), water molecules could well penetrate into the micropores of BCPs and form water clusters via intermolecular hydrogen bonding. The sorbing ability order of the different types of BCPs varied depending on the examined RH. At 33% RH, the k_2 roughly followed a decreasing order of soot ($0.5\text{--}5.25 \times 10^{-5}$ g $\text{mg}^{-1}\text{s}^{-1}$) > woody BCPs ($1.57\text{--}1.90 \times 10^{-5}$ g $\text{mg}^{-1}\text{s}^{-1}$) > herbal BCPs ($0.34\text{--}2.07 \times 10^{-5}$ g $\text{mg}^{-1}\text{s}^{-1}$); however, no clear trend was shown for high humidity conditions (e.g., RH = 94%), mainly resulting from the larger variances within the group of herbal BCPs and soot group.

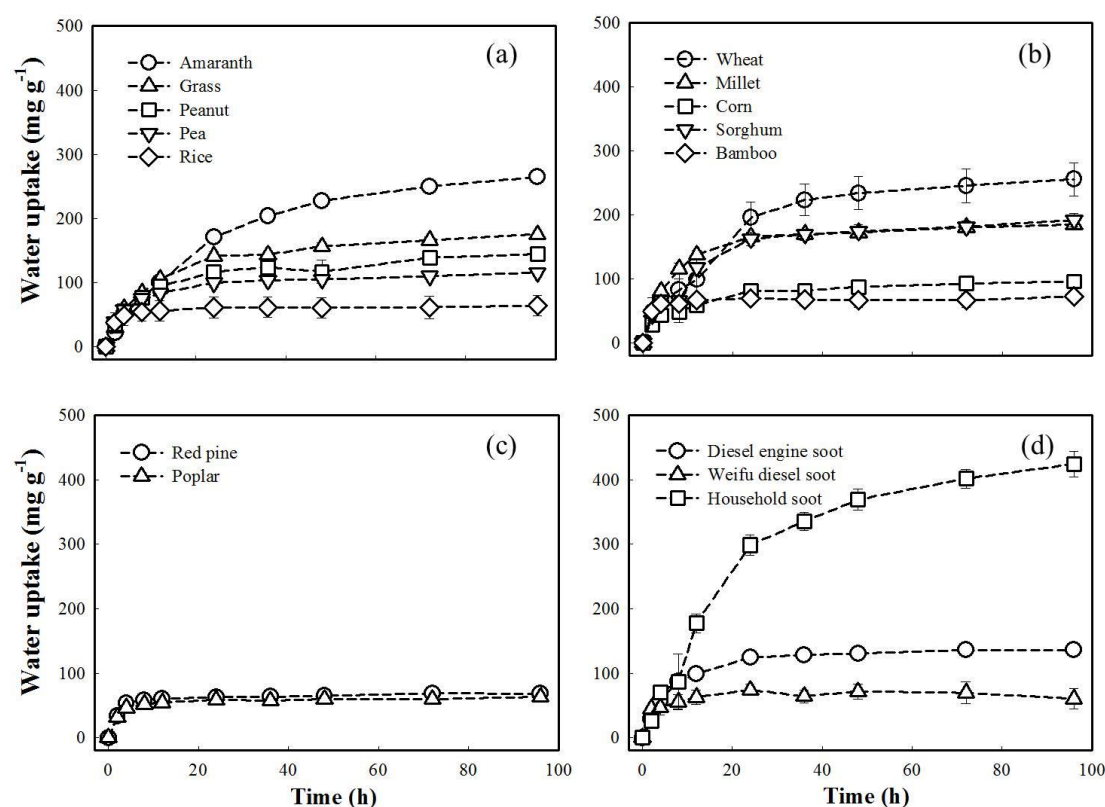


Figure 4. Sorption kinetics of water vapor plotted as water uptake (mg g^{-1}) vs. time (h) at 94% relative humidity. (a) Subgroup 1 of herbal BCPs. (b) Subgroup 2 of herbal BCPs. (c) Woody BCPs. (d) Soot. Error bars represent standard variations from

duplicate samples.

Like for the equilibrium water uptake, the relationships were built between the k_2 and the contents of OC_{TGA}, OC_{AE}, EC, dissolved minerals, major ionic species (NH₄⁺, Cl⁻, C₂O₄²⁻ and SO₄²⁻), and total porosity, respectively at 94% RH (Figure 5). The regression relationships at 33% RH were presented in Figure S11. The regression accuracy (R^2 and P) values at 33%, 47%, and 94% RH are summarized in Table S11. At 94% RH, a strong positive correlation was observed between the k_2 and the total porosity ($R^2 = 0.82$, $P < 0.0001$), likely resulting from the pore-filling/condensation mechanism. A similar mechanism has been previously proposed to account for the positive correlation observed between the water vapor sorption kinetics and the porosity of activated carbon under high humidity conditions (Nakamura et al., 2010; Velasco et al., 2016). Likewise, a positive correlation ($R^2 = 0.45$, $P = 0.0061$) was observed between the k_2 and the EC content, which could be ascribed to the formation of rigid pore structures dominantly by graphitized carbons (EC) (Han et al., 2014). The correlation between k_2 and total porosity observed at 94% RH disappeared under low and medium humidity conditions. At 33% RH, relatively good positive correlations were observed with the contents of OC_{TGA} ($R^2 = 0.47$, $P = 0.0046$), OC_{AE} ($R^2 = 0.44$, $P = 0.007$), NH₄⁺ ($R^2 = 0.77$, $P < 0.0001$), and Cl⁻ ($R^2 = 0.60$, $P = 0.0007$), but not with SO₄²⁻ ($R^2 = 0.11$, $P = 0.2286$) or dissolved minerals ($R^2 = 0.08$, $P = 0.31$). The positive correlations with these constituents were not shown under medium and high humidity conditions. Thus, the constituents of OC and NH₄⁺- and Cl⁻-salts likely provided the primary high affinity, active sites for sorption of water vapor under low humidity conditions.

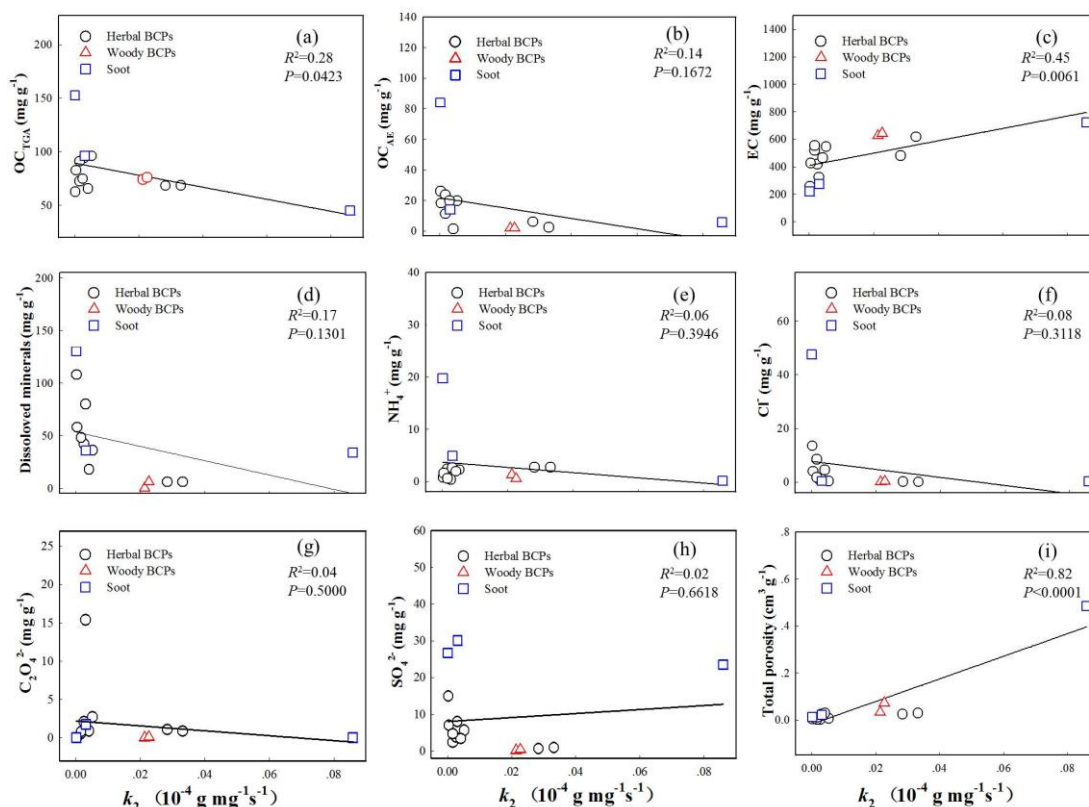


Figure 5. Relationships between pseudo-second water uptake rate constant (k_2) ($g\ mg^{-1}\ s^{-1}$) vs. compositional and pore property parameters for the group of BCPs at 94% relative humidity. (a) TGA-measured organic carbon (OC_{TGA}). (b) Alkali-extracted organic carbon (OC_{AE}). (c) Elemental carbon (EC). (d) Dissolved minerals. (e) Ammonium (NH_4^+). (f) Chloride (Cl^-). (g) Oxalate ($C_2O_4^{2-}$). (h) Sulfate (SO_4^{2-}). (i) Total porosity.

4. Conclusion

The hygroscopic properties of 15 different BCPs (herbal, woody, and soot) were systematically investigated using gravimetric method and DRIFTS. The mechanisms and factors controlling the equilibrium and kinetic water uptake differed among the types of BCPs and depended heavily on the humidity conditions. Linear correlation analyses indicated that the equilibrium water uptake by the tested group of BCPs positively correlated to the contents of OC (OC_{TGA} and OC_{AE}), dissolved minerals,

and NH_4^+ - and Cl^- -salts under high humidity conditions, and weakly to the contents of OC only under low humidity conditions. By contrast, negative correlations were observed between the equilibrium water uptake and the EC content. The low water uptake by the woody BCPs could be attributed to the very low contents of OC and dissolved minerals. Thus, the equilibrium water uptake by BCPs was mainly controlled by the hygroscopic constituents of OC and dissolved minerals/salts. The kinetic water uptake by the BCPs could be well described by the pseudo-second-order kinetic model. The calculated rate constant (k_2) positively correlated to the contents of OC_{TGA} , OC_{AE} , and NH_4^+ - and Cl^- -salts under low humidity conditions, and to the total porosity only under high humidity conditions. The fast water uptake kinetics under low humidity conditions was attributed to the binding to high affinity, active sites (OC and salts), which could sorb water vapor fast but had limited contents. Alternatively, the slow water uptake kinetics under high humidity conditions was attributed to pore-filling/condensation of water molecules within the micro- and mesopores of BCPs. This study highlights that the hygroscopic properties of BCPs rely on compositional and structural properties of BCPs as well as humidity conditions.

Author contributions. DZ provided the original idea and prepared the paper with contributions from all co-authors. MW and YC designed and conducted the research, HF, XQ were involved in the development of the analysis methods. BL, ST reviewed the written document.

Competing interests. The authors declare that they have no conflict of interest.

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Appendix A.

Detailed characterization results of the different BCPs can be found in Tables

S1-S7. Table S8 lists accuracy (R^2 and P) values for regression on equilibrium water uptake against different variables. Tables S9-S10 present pseudo-first/second-order kinetic model fitting parameters. Table S11 lists accuracy (R^2 and P) values for regression on k_2 against different variables. Figure S1-S4 displays spectroscopic characterization of different BCPs. Figure S5 displays sorption isotherms at selected humidity obtained by using saturated aqueous salt solutions. Figure S6 compares equilibrium water uptake measured by the two different gravimetric methods. Figure S7 displays relationships between equilibrium water uptake and different variables at 23% RH. Figure S8 displays the relationship between measured values of equilibrium water uptake at 94% relative humidity vs. predicted values obtained by binary factor regression based on contents of OC_{TGA} and dissolved minerals for the group of BCPs. Figure S9 compares equilibrium water uptake measured by DRIFTS and gravimetric method at high RH. Figure S10 displays sorption kinetics of water uptake at 33% RH. Figure S11 displays relationships between k_2 and different variables at 33% RH.

References

- Adapa, S., Swamy, D. R., Kancharla, S., Pradhan, S., and Malani, A.: Role of mono- and divalent surface cations on the structure and adsorption behavior of water on mica surface, *Langmuir*, 34 (48), 2018.
- Andreae, M. O., and Gelencser, A.: Black carbon or brown carbon? The nature of light-absorbing carbonaceous aerosols, *Atmos. Chem. Phys.*, 6, 3131–3148, 2006.
- Archango, B. S., Araujo, J. R., Silva, A. M., Capaz, R. B., Falcao, N. P. S., Jorio, A., and Achete, C. A.: Chemical analysis and molecular models for calcium-oxygen-carbon interactions in black carbon found in fertile Amazonian anthrosoils, *Environ. Sci. Technol.*, 48, 7445–7452, 2014.
- Bond, T. C., Doherty, S. J., Fahey, D. W., Forster, P. M., Berntsen, T., DeAngelo, B. J., Flanner, M. G., Ghan, S., Kärcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P. K., Sarofim, M. C., Schultz, M. G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S., Bellouin, N., Guttikunda, S. K., Hopke, P. K., Jacobson, M. Z., Kaiser, J. W., Klimont, Z., Lohmann, U., Schwarz, J. P., Shindell, D., Storelvmo, T., Warren, S. G., and Zender, C. S.: Bounding the role of black carbon in the climate system: A scientific assessment, *J. Geophys. Res. Atmos.*, 118, 5380–5552, 2013.
- Buchholz, A., and Mentel, T. F.: Hygroscopic growth of oxalic acid and its ammonium, sodium and potassium salts, *Geophys. Res. Abs.*, 10, EGU2008-A-03795.
- Cancado, L. G., Takai, K., Enoki, T., Endo, M., Kim, Y. A., and Mizusaki, H.: General equation for the determination of the crystallite size L_a of nanographite by Raman spectroscopy, *Appl. Phys. Lett.*, 88 (16), 2006.
- Carrico, C. M., Petters, M. D., Kreidenweis, S. M., Sullivan, A. P., McMeeking, G. R., Levin, E. J. T., Engling, G., Malm, W. C., and Collett Jr. J. L.: Water uptake and chemical composition of fresh aerosols generated in open burning of biomass,

588 Atmos. Chem. Phys., 10, 5165–5178, 2010.

589 Day, D. E., Hand, J. L., Carrico, C. M., Engling, G., and Malm, W. C.: Humidification
 590 factors from laboratory studies of fresh smoke from biomass fuels, J. Geophys. Res.,
 591 111, D22202, 2006.

592 Fletcher, A. J., Uygur, Y., and Mark Thomas, K.: Role of surface functional groups in
 593 the adsorption kinetics of water vapor on microporous activated carbons, J. Phys.
 594 Chem. C., 111, 8349–8359, 2007.

595 Freney, E. J., Martin, S. T., and Buseck, P. R.: Deliquescence measurements of
 596 potassium salts, American Geophysical Union, Fall Meeting, 2007.

597 Ghorai, S., Laskin, A., and Tivanski, A. V.: Spectroscopic evidence of keto-enol
 598 tautomerism in deliquesced malonic acid particles, J. Phys. Chem. A., 115,
 599 4373–4380, 2011.

600 Gu, W., Li, Y., Zhu, J., Jia, X., Lin, Q., Zhang, G., Ding, X., Song, W., Bi, X., Wang,
 601 X., and Tang, M.: Investigation of water adsorption and hygroscopicity of
 602 atmospherically relevant particles using a commercial vapor sorption analyzer,
 603 Atmos. Meas. Tech., 10, 3821–3832, 2017.

604 Gustafsson, R. J., Orlov, A., Badger, C. L., Griffiths, P. T., Cox, R. A., and Lambert, R.
 605 M.: A comprehensive evaluation of water uptake on atmospherically relevant
 606 mineral surfaces: DRIFT spectroscopy, thermogravimetric analysis and aerosol
 607 growth measurements, Atmos. Chem. Phys., 5, 3415–3421, 2005.

608 Haddrell, A. E., Davies, J. F., and Reid, J. P.: Dynamics of particle size on inhalation
 609 of environmental aerosol and impact on deposition fraction, Environ. Sci. Technol.,
 610 49, 14512–14521, 2015.

611 Han, C., Liu, Y., and He, H.: Role of organic carbon in heterogeneous reaction of NO₂
 612 with soot, Environ. Sci. Technol., 47, 3174–3181, 2013.

613 Han, L. F., Sun, K., Jin, J., Wei, X., Xia, X. H., Wu, F. C., Gao, B., and Xing, B. S.:
614 Role of structure and microporosity in phenanthrene sorption by natural and
615 engineered organic matter, *Environ. Sci. Technol.*, 48, 11227–11234, 2014.

616 Jing, B., Peng, C., Wang, Y., Liu, Q., Tong, S., Zhang, Y., and Ge, M.: Hygroscopic
617 properties of potassium chloride and its internal mixtures with organic compounds
618 relevant to biomass burning aerosol particles, *Sci. Rep.*, 2, 43573, 2017.

619 Keiluweit, M., Nico, P. S., Johnson, M. G., Kleber, M.: Dynamic molecular structure
620 of plant biomass-derived black carbon (biochar), *Environ. Sci. Technol.*, 44 (4),
621 1247–1253, 2010.

622 Lewis, K. A., Arnott, W. P., Moosmuller, H., Chakrabarty, R. K., Carrico, C. M. ,
623 Kreidenweis, S. M., Day, D. E., Malm, W. C., Laskin, A., Jimenez, J. L., Ulbrich, I.
624 M., Huffman, J. A., Onasch, T. B., Trimborn, A., Liu, L., and Mishchenko, M. I.:
625 Reduction in biomass burning aerosol light absorption upon humidification: roles
626 of inorganically-induced hygroscopicity, particle collapse, and photoacoustic heat
627 and mass transfer, *Atmos. Chem. Phys.*, 9, 8949–8966, 2009.

628 Li, J., Posfai, M., Hobbs, P. V., and Buseck, P. R.: Individual aerosol particles from
629 biomass burning in southern Africa: 2. Compositions and aging of inorganic
630 particles, *J. Geophys. Res.*, 108, 8484, 2003.

631 Lian, F., and Xing, B.: Black carbon (biochar) in water/soil environments: Molecular
632 structure, sorption, stability, and potential risk, *Environ. Sci. Technol.*, 51,
633 13517–13532, 2017.

634 Liu, P., Song, M., Zhao, T., Gunthe, S. S., Ham, S., He, Y., Qin, Y. M., Gong, Z.,
635 Amorim, J. C., Bertram, A. K., and Martin, S. T.: Resolving the mechanisms of
636 hygroscopic growth and cloud condensation nuclei activity for organic particulate
637 matter, *Nat. Commun.*, 9, 4076, 2018.

638 Ma, Q., He, H., Liu, Y.: In situ DRIFTS study of hygroscopic behavior of mineral
639 aerosol, *J. Environ. Sci.*, 22(4), 555–560, 2010.

640 Matthews, H. D., Gillett, N. P., Stott, P. A., and Zickfeld, K.: The proportionality of
641 global warming to cumulative carbon emissions, *Nature*, 459, 829–832, 2009.

642 Nakamura, M., Ohba, T., Branton, P., Kanoh, H., and Kaneko, K.: Equilibrium-time
643 and pore width dependent hysteresis of water adsorption isotherm on hydrophobic
644 microporous carbons, *Carbon*, 48, 2549–2553, 2010.

645 Nguyen, T. X., and Bhatia, S. K.: How water adsorbs in hydrophobic nanospaces, *J.*
646 *Phys. Chem. C*, 115, 16606–16612, 2011.

647 Nienow, A. M., Roberts, J. T.: Heterogeneous chemistry of carbon aerosols, *Rev. Phys.*
648 *Chem.*, 57, 105–128, 2006.

649 Ohba, T., and Kaneko, K.: Kinetically forbidden transformations of water molecular
650 assemblies in hydrophobic micropores, *Langmuir*, 27, 7609–7613, 2011.

651 Petters, S. S., Pagonis, D., Claflin, M. S., Levin, E. J. T., Petters, M. D., Ziemann, P. J.,
652 and Kreidenweis, S. M.: Hygroscopicity of organic compounds as a function of
653 carbon chain length and carboxyl, hydroperoxy, and carbonyl functional groups, *J.*
654 *Phys. Chem. A*, 121 (27), 5164–5174, 2017.

655 Pimenta, M. A., Dresselhaus, G., Dresselhaus, M. S., Cancado, L. G., Jorio, A., and
656 Saito, R.: Studying disorder in graphite-based systems by Raman spectroscopy,
657 *Phys. Chem. Chem. Phys.*, 9, 1276–1291, 2007.

658 Posfai, M., Simonics, R., Li, J., Hobbs, P. V., and Buseck, P. R.: Individual aerosol
659 particles from biomass burning from southern Africa: 1. Compositions and size
660 distributions of carbonaceous particles, *J. Geophys. Res.*, 108, 8483, 2003.

661 Powelson, M. H., Espelien, B. M., Hawkins, L. N., Galloway, M. M., and De Haan, D.
662 O.: Brown carbon formation by aqueous-phase carbonyl compound reactions with

663 amines and ammonium sulfate, *Environ. Sci. Technol.*, 48, 985–993, 2014.
 664 Preturlan, J. G. D., Vieille, L., Quilgotti, S., and Favergeon, L.: Comprehensive
 665 thermodynamic study of the calcium sulfate–water vapor system. Part 1:
 666 experimental measurements and phase equilibria, *Ind. Eng. Chem. Res.*, 58, 22,
 667 9596–9606, 2019.
 668 Qiu, C., Khalizov, A. F., and Zhang, R.: Soot aging from OH-Initiated oxidation of
 669 toluene, *Environ. Sci. Technol.*, 46, 9464–9472, 2012.
 670 Reid, J. S., Koppmann, R., Eck, T. F., and Eleuterio, D. P.: A review of biomass
 671 burning emissions part II: intensive physical properties of biomass burning particles,
 672 *Atmos. Chem. Phys.*, 5, 799–825, 2005.
 673 Ribeiro, A. M., Sauer, T. P., Grande, C. A., Moreira, R. F. P. M., Loureiro, J. M., and
 674 Rodrigues, A. E.: Adsorption equilibrium and kinetics of water vapor on different
 675 adsorbents, *Ind. Eng. Chem. Res.*, 47, 18, 7019–7026, 2008.
 676 Rosas, J. M., Bedia, J., Miraso, J. R., and Cordero, T.: Preparation of hemp-derived
 677 activated carbon monoliths, adsorption of water vapor, *Ind. Eng. Chem. Res.*, 47, 4,
 678 1288–1296, 2008.
 679 Schwarz, J. P., Spackman, J. R., Fahey, D. W., Gao, R. S., Lohmann, U., Stier, P.,
 680 Watts, L. A., Thomson, D. S., Lack, D. A., Pfister, L., Mahoney, M. J.,
 681 Baumgardner, D., Wilson, J. C., and Reeves, J. M.: Coatings and their enhancement
 682 of black carbon light absorption in the tropical atmosphere, *J. Geophys. Res.*, 113,
 683 D03203, 2008.
 684 Seisel, S., Pashkova, A., Lian, Y., and Zellner, R.: Water uptake on mineral dust and
 685 soot: a fundamental view of the hydrophilicity of atmospheric particles, *Faraday*
 686 *Discuss.*, 130, 437–451, 2005.
 687 Sjogren, S., Gysel, M., Weingartner, E., Baltensperger, U., Cubison, M. J., Coe, H.,

688 Zardini, A. A., Marcolli, C., Krieger, U. K., and Peter, T.: Hygroscopic growth and
 689 water uptake kinetics of two-phase aerosol particles consisting of ammonium
 690 sulfate, adipic and humic acid mixtures, *Aerosol Science*, 38, 157–171, 2007.

691 Song, J., Peng, P., and Huang, W.: Black carbon and kerogen in soils and sediments. 1.
 692 Quantification and characterization, *Environ. Sci. Technol.*, 36, 3960–3967, 2002.

693 Song, W., and Boily, J. F.: Water vapor adsorption on goethite, *Environ. Sci. Technol.*,
 694 47, 7171–7177, 2013.

695 Stanislav, V., Vassilev, D. B., Andersen, L. K., and Vassileva, C. G.: An overview of
 696 the composition and application of biomass ash. Part 1. Phase-mineral and chemical
 697 composition and classification, *Fuel*, 105, 40–76, 2013.

698 Stemmler, K., Vlasenko, A., Guimbaud, C., and Ammann, M.: The effect of fatty acid
 699 surfactants on the uptake of nitric acid to deliquesced NaCl aerosol, *Atmos. Chem.*
 700 *Phys.*, 2008, 8 (17), 5127–5141, 2008.

701 Suda, S. R., Petters, M. D., Yeh, G. K. Strollo, C., Matsunaga, A., Faulhaber, A.,
 702 Ziemann, P. J., Prenni, A. J., Carrico, C. M., Sullivan, R. C., and Kreidenweis, S.
 703 M.: Influence of functional groups on organic aerosol cloud condensation nucleus
 704 activity, *Environ. Sci. Technol.*, 48, 10182–10190, 2014.

705 Velasco, L. F., Remy Guillet, N. R., Dobos, G., Thommes, M., and Lodewyckx, P.:
 706 Towards a better understanding of water adsorption hysteresis in activated carbons
 707 by scanning isotherms, *Carbon*, 96, 753–758, 2016.

708 Wei, S., Zhu, M., Fan, X., Song, J., Peng, P., Li, K., Jia, W., and Song, H.: Influence
 709 of pyrolysis temperature and feedstock on carbon fractions of biochar produced
 710 from pyrolysis of rice straw, pine wood, pig manure and sewage sludge,
 711 *Chemosphere*, 218, 624–631, 2019.

712 Xiao, J., Liu, Z., Kim, K., Chen, Y., Yan, J., Li, Z., and Wang, W.: S/O-Functionalities

713 on modified carbon materials governing adsorption of water vapor, *J. Phys. Chem.*
 714 *C.*, 117, 23057–23065, 2013.

715 Xiao, X., Chen, B., Chen, Z., Zhu, L., and Schnoor, J. L.: Insight into multiple and
 716 multilevel structures of biochars and their potential environmental applications: A
 717 critical review, *Environ. Sci. Technol.*, 52(9), 5027–5047, 2018.

718 Yuan, J. F., Huang, X. F., Cao, L. M., Cui, J., Zhu, Q., Huang, C. N., Lan, Z. J., and
 719 He, L.Y.: Light absorption of brown carbon aerosol in the PRD region of China,
 720 *Atmos. Chem. Phys. Discuss.*, 15, 28453–28482, 2015.

721 Yuan, W., Li, X., Pan, Z., Connell, L. D., Li, S., and He, J.: Experimental
 722 investigation of interactions between water and a lower silurian chinese shale,
 723 *Energy Fuels*, 28, 4925–4933, 2014.

724 Zhang, R., Khalizov, A., Wang, L., Hu. M., and Xu, W.: Nucleation and growth of
 725 nanoparticles in the atmosphere, *Chem. Rev.*, 112, 3, 1957–2011, 2012.

726 Zuend, A., Marcolli, C., Booth, A. M., Lienhard, D. M., Soonsin, V., Krieger, U. K.,
 727 Topping, D. O., McFiggans, G., Peter, T., and Seinfeld, J. H.: New and extended
 728 parameterization of the thermodynamic model AIOMFAC: Calculation of activity
 729 coefficients for organic-inorganic mixtures containing carboxyl, hydroxyl, carbonyl,
 730 ether, ester, alkenyl, alkyl, and aromatic functional groups, *Atmos. Chem. Phys.*, 11,
 731 9155–9206, 2011.