Deposition of light-absorbing particles in glacier snow of the Sunderdhunga Valley, the southern forefront of Central Himalaya

Jonas Svensson[1,2], Johan Ström[3], Henri Honkanen[4], Eija Asmi[1], Nathaniel B. Dkhar[5], Shresth Tayal[5,6], Ved P. Sharma[5,6], Rakesh Hooda[1], Matti Leppäranta[4], Hans-Werner Jacobi[2], Heikki Lihavainen[7,1], Antti Hyvärinen[1]

1 Atmospheric Composition Research, Finnish Meteorological Institute, Helsinki, Finland

2 Université Grenoble Alpes, CNRS, IRD, INP-G, IGE, Grenoble, France

3 Department of Environmental Science, Stockholm University, Stockholm, Sweden

4 Institute for Atmospheric and Earth System Research, Faculty of Science, University of Helsinki, Helsinki, Finland

5 The Energy and Resource Institute, (TERI), New Delhi, India

6 TERI School of Advanced Studies (TERI SAS), New Delhi, India

7 Svalbard Integrated Arctic Earth Observing System, Longyearbyen, Norway

Correspondence to: jonas.svensson@fmi.fi

Abstract

Anthropogenic activities on the Indo-Gangetic Plain emit vast amounts of light-absorbing particles (LAP) into the atmosphere, modifying the atmospheric radiation ~~scheme~~ state. With transport to the nearby Himalayan mountains and deposition to its surfaces the particles contribute to glacier and snowmelt via darkening of the highly reflective snow. The Central Himalayas have been identified as a region where LAP are especially pronounced in glacier snow, but still remain a region where measurements of LAP in the snow are scarce. Here we study the deposition of LAP in five snow pits sampled in 2016 (and one from 2015) within one km from each other from two glaciers in the Sunderdhunga valley, state of Uttarakhand, India, Central Himalaya. The snow pits display a distinct ~~melt~~ enriched LAP layer interleaved by younger snow above, and older snow below. The LAP exhibit a distinct vertical distribution in these different snow layers. For the analyzed elemental carbon (EC), the younger snow layers in the different pits show similarities, ~~and~~ which can be characterized by a deposition constant of about 50 µg m^{-2} per mm^{-1} snow water equivalent (SWE) while the old snow layers also indicate similar values, ~~and can be~~ described ~~with~~ by a deposition constant of roughly 150 µg m^{-2} per mm^{-1} SWE. The ~~melt~~ enriched LAP layer, contrarily, display no similar trends between the pits. Instead, it is characterized by very high amounts of LAP, and differ in orders of magnitude for concentration between the pits. The ~~melt~~ enriched LAP layer is likely a result of strong melting that took place during the summers of 2015 and 2016, as well as possible lateral transport of LAP. The

mineral dust fractional absorption is slightly below 50 % for the young and old snow layer, whereas it~~n the melt layer~~ is the dominating light absorbing constituent in the enriched LAP layer, thus, highlighting the importance of dust in the region. Our results indicate the problems with complex topography in the Himalaya, but nonetheless, can be useful in large-scale assessments of LAP in Himalayan snow.

1 Introduction

Aerosol particles in the Indo-Gangetic Plain (IGP) are produced in great mass and number. Being especially prominent in the pre-monsoon season, a large fraction of the airborne aerosols are carbonaceous particles, consisting of organic carbon (OC) and black carbon (BC). Originating from the combustion of fossil fuels and biomass, the particles form the atmospheric brown cloud—known to modify the atmospheric radiation ~~scheme~~ state (Lau et al., 2006; Menon et al., 2010; Ramanathan and Carmichael, 2008). Through air mass transport the aerosol can be conveyed and lifted from the IGP to its northern barrier, the mountains of Himalaya (e.g. Hooda et al., 2018; Kopacz et al., 2011; Raatikainen et al., 2014; Zhang et al., 2015). Covered with vast amounts of snow and ice, the Himalayan cryosphere is affected by the deposition of carbonaceous aerosol onto its surface (e.g. He et al., 2018; Jacobi et al., 2015; Ménégoz et al., 2014; Xu et al., 2009). This is due to the particulates and especially BC effectiveness in reducing the snow albedo (Warren and Wiscombe, 1980), which ultimately leads to accelerated snow melt (Flanner et al., 2007; Jacobi et al., 2015; Jacobson, 2004; Ming et al., 2012).

In addition to BC and OC, other particles such as mineral dust (MD) and snow microbes (collectively known as light-absorbing particles (LAP)) are also of importance in reducing snow albedo (e.g. Skiles et al. 2018). In Himalayan snow and ice, the LAP content has been shown to vary significantly, both spatially and temporally (e.g. see review by Gertler et al., 2016). Further, an extensive compilation of BC measurements in snow over the Tibetan Plateau is presented in the supplement of He et al. (2018), with concentrations ranging from 1 to 3600 ppb_w in the region termed as Himalaya. In addition to long range transported LAP, local sources within the Tibetan plateau have also been documented to be significant in some regions (e.g. Li et al., 2016), creating several different sources of LAP in the snow. Varying meteorology and terrain induced exchange processes (advection and turbulence) in the mountains further complicates the interplay between the atmospheric deposition of LAP and the snow surfaces.

~~Observations are further supported by modeling studies, which indicate~~ Recent modeling studies have reported analogous results, indicating certain sub-regions of the Himalaya to be especially vulnerable to LAP deposition. Santra et al. (2019) ~~recently~~ simulated the BC impact on snow albedo and glacier runoff in the Hindu Kush-Himalaya region. The authors identified a hot-spot zone for BC in the vicinity of Manora peak, located in the Indian state of Uttarakhand, central Himalaya (also sometimes called western Himalaya depending on classification). The BC induced a greater albedo reduction on glacier snow in the vicinity of this hot spot area compared to other areas in the Hindu Kush-Himalayan area. Similarly, another modeling study simulated the impact of LAP on High Mountain Asia snow albedo and its associated forcing and identified the same general area as a region where snow is especially affected by LAP-caused snow darkening (Sarangi et al., 2019). Both of these studies (as well as the work of He et al., 2018) emphasized the need for more *in situ* measurements of LAP in the snow of this region of the Himalaya.

Previously, we reported in Svensson et al. (2018) the measured LAP concentrations and properties in the snow from two glaciers in the Sunderdhunga valley, located in Uttarakhand, India, central Himalaya. While we mainly focused on the surface snow layer and characterizing the LAP, results from one 1.2 m deep snow pit were also presented. Based on the LAP concentration profile and pit stratigraphy, the pit was estimated to represent 5 seasons. Newly sampled snow pits have since then been analyzed from the same two glaciers, along with available automatic weather station (AWS) data from the same valley. Here we revisit the previous interpretation of the published pit (in Svensson et al., 2018), and report the results of our newly sampled snow pits. By comparing the BC profiles among 6 pits we aim at quantifying the deposition of elemental carbon (EC; used here as a proxy for BC) in this area of the Himalaya. In addition, we explore the relative contribution of MD to LAP in the different pits.

2 Methodology

2.1 Glaciers snow sampling and filtration

Snow was collected on Bhanolti and Durga Kot glaciers during a field campaign in the Sunderdhunga valley (located in the Bageshwar district) in October of 2016. The two glaciers are positioned adjacent to each other in a general northeast-southwest orientation (cf. Fig. 1) on the southern fringe of the Himalayan mountain range and are further described in Svensson et al. (2018). Local emissions of carbonaceous aerosol in the Sunderdhunga valley are very limited. The valley is not accessible by car and the glaciers are at a three to four-day hike from the nearest road. On route to the glaciers the last settlement is Jatoli, located in a river valley at an elevation of 2400 m. a.s.l. about 10 km southeast in a perpendicular orientation to the glacier valley. Biomass burning is a common practice for cooking and heating in Jatoli, thus some emissions from the village may enter the glacier valley. It is expected, however, that the majority of carbonaceous particles in the glacier valley originates from regional and long-distance transport. The relatively low elevation span as well as the glaciers' position on the southern slopes of the Himalayan mountains nonetheless, make them more prone to LAP deposition compared to other glaciers in the Himalaya and Tibetan plateau. Previous studies have reported elevated LAP content in lower elevation snow for Himalayan glaciers (e.g. Ming et al., 2013), and higher concentrations of LAPs in glaciers on the southern edge of the Himalaya (e.g. Xu et al., 2009).

On Durga Kot glacier two snow pits (hereafter Pit A and B; Fig. 1) were dug in the vicinity of each other (~20 m) in an reachable area of the percolation zone of the glacier. Bhanolti glacier was more easily accessible, and the three excavated snow pits (hereafter Pit C, D, E; Fig. 1) were spread out over a greater distance (~500 m) on the glacier (see table 1 and Fig. 1 for additional information). The depth of the pits depended on the level at which a hard layer was found, and digging could not be further conducted with the reinforced shovels with a sharpened edge. The deepest snow pit that was analyzed previously in Svensson et al. (2018), referred to as pit 5 in that study, is from Bhanolti glacier in September of 2015, and we denote as Pit F in the subsequent sections of this manuscript. As for the other pits from 2016, the depth of Pit F was governed by the depth at which the hard layer was encountered.

Three distinctly different colored snow layers could be observed repeating in all but one of the year 2016 pits: a relatively thin (on the order of centimeters) very dark layer was ~~separated by~~ wedged in-between white snow above and more grey appearing snow below (See for e.g. pits B and D in Fig. S1a-b). Due to this stratigraphy, we hereafter simply refer to the whitest snow as young snow, the darkest layer as the ~~melt~~ enriched LAP layer, and the grey snow as old snow. Representative samples ranging from 3 to 10 cm thick layers were taken throughout each pit for analysis of LAP. Snow density measurements were conducted with a snow density kit in the upper part of the pits (in 5 cm increments) by weighing the known volume of the sampler filled with snow. The observed densities ranged between 0.29 and 0.46 g cm^{-3} (see table 1 for details). Density measurements were not possible below the ~~melt~~ enriched LAP layer due to the hard snow. For these layers the density was assumed 0.5 g cm^{-3} (to represent aged snow) in our further analyzes. Snow density measurements were not conducted for Pit F, and we assigned a density of 0.35 g cm^{-3} for the top layer (0-3 cm; similar to observations made in 2016), followed by 0.4 g cm^{-3} between 3-10 cm depth, and 0.5 g cm^{-3} for all layers below 10 cm. Since the snow samples could not be transported in a solid phase back to the laboratory, they were melted and filtered at the nearby base camp using the same principles as in Svensson et al. (2018). Filters were transported back to the analysis laboratory in petri slides.

2.2 Meteorological observations

In September 2015 an AWS was installed next to the glacier ablation zone of Durga Kot (Fig. 1) about 1.5 km northwards at an elevation below the snow sampling sites. The AWS is equipped with instruments for air temperature, relative humidity (HC2S3-L Temperature and relative humidity probe manufactured by Rotronic, with 41303-5A Radiation shield), shortwave (SW) and longwave (LW) radiation (upward and downward) (CNR4 Four-component net radiometer manufactured by Kipp & Zonen), wind speed and direction (05103-L Wind monitor manufactured by R. M. Young), and snow depth (Campbell Scientific SR50A-L Ultrasonic Distance Sensor). In this paper we use the snow depth data between September 2015 and September 2017 to estimate the local precipitation. The original snow depth data, logged once every 10 minutes was filtered to daily resolution by applying a moving median window of 24 hours and for the noon value of each day in further analyzes. This filtering removed much of the signal noise. However, before this filtering was applied the data was reduced using several logical conditions such as: the incoming SW radiation is greater than outgoing SW radiation (to remove errors due to sensors covered by snow), and the surface albedo is greater than 0.2 (to ensure snow cover as the ground albedo was measured at 0.17). Finally, the consistency between the daily albedo and snow depth was inspected using data presented in Figure S~~1~~2a. Each day the snow depth increased was interpreted as precipitation, and to arrive at an estimate of the snow water equivalent (SWE), the fresh snow density is assumed to be 100 kg m^{-3} (Helfricht, et al., 2018). The solid precipitation derived based on the cumulative SWE is presented in Figure S~~1~~2b.

2.3 Filter analysis

The analysis of filters followed the procedure in Svensson et al. (2018), with transmission measurements coupled with thermal-optical analysis. According to the measurement nomenclature (Petzold et al., 2013), the carbonaceous constituents measured are EC and OC. The measurement method briefly follows the procedure of placing a filter punch in a custom-built particle soot absorption photometer (PSAP) to measure the transmittance (at λ = 526 nm; Krecl et al., 2007)—providing an optical depth for all of the particles captured by the filter. The filter punch is then placed in an OCEC analyzer (Sunset instrument, using the EUSAAR_2 protocol) to determine the OC and EC mass, followed by another measurement with the PSAP. The OCEC analysis removes the carbonaceous species and, thus, by comparing the PSAP results obtained before and after the analysis, the relative contribution of the light absorption by EC particles in the total particles optical depth is obtained. The remaining optical depth we attribute as non-EC material. This fraction of the total optical thickness we report as the percentage of the mineral dust absorption on the filter samples (expressed as f_D). For further details concerning the measurements see Svensson et al (2018).

Some of the filter samples (N=17, out of 91) were saturated with too much light absorbing material prohibiting reliable EC measurements despite reducing the sample to a melted equivalent of only 30 mL. To mitigate this problem, we calculated the EC indirectly from the analyzed total carbon (TC) for the saturated samples. From OCEC analysis TC is the most robust measured constituent, since it includes both OC and EC and is not affected by their split point, which may be incorrectly placed for very dark filters (Chow et al., 2001). A slope of 0.099 for the EC:TC ratio for filter samples considered non-saturated ~~A linear relation was fitted for non-saturated filter samples and the obtained correlation of $EC = 0.10TC + 0.12$~~ was used to reconstruct the EC content for the filter samples containing high amounts of absorbing particles (see details in supplement and Fig. S~~3~~3a-b). The slope compares well with the slopes reported for air samples collected at two sites in the Himalayas about 550 km south-east from Sunderdhunga in the Kathmandu valley 32 km (altitude of 2150 m a.s.l.) east of Kathmandu, and Langtang 60 km north of Kathmandu (altitude of 3920 m a.s.l) (Caricco et al., 2003). There, the authors found that the EC/TC ratio was 0.17 for both sites during the summer monsoon season, but between

0.10 and 0.13 during what they described as the ramp-up period and the peak concentration season. The snow samples do not have an upper limit for particles sizes, whereas the air samples were collected as PM2.5 (particulate matter collected below an aerodynamic diameter of 2.5 µm). The slopes are rather similar to our value, and the authors found as well a very strong correlation of 0.89 (r^2) between monthly average EC and OC.

3 Results and discussion

3.1 EC deposition in young and old snow samples

When the EC content is analyzed from filtered snow samples, a common practice is to convert the results into mass concentrations [EC], given per volume or mass of melt water (e.g. µg L^{-1} or ng g^{-1}). A spread in results is often largely due to local processes and specific sampling layer thicknesses. The mass deposition per unit area $\widetilde{EC}$, on the other hand, can be expected to be less variable with increasing number of layers used to calculate this value. The deposition in each layer is calculated according to:

$$\widetilde{EC_i} = [EC]_i \frac{\rho_{s_i}}{\rho_w} d_i \quad (1)$$

were ρ_s and ρ_w are snow and liquid water densities, respectively. The index i, is the number of the sampled layer from top to bottom, and $\rho_s/\rho_w\ d$ is the SWE thickness, d_{SWE}. The $\widetilde{EC_i}$ and d_{SWEi} are transformed to cumulative plots by integrating over the layers from the surface to the bottom. These profiles are presented in Fig. 2a-f (with each sampling layer represented by a square).

The visible snow pit stratigraphy described above in section 2.1 can be observed in the pit profiles. At the top, the accumulated EC ($ECacc$) as a function of the accumulated d_{SWE} ($SWEacc$) portray the young snow layers, whereas in the bottom of the pits the data points represent the old snow layers (Fig. 2a-f). This pattern (with both young and old snow layers) is visible in pits A, B, C, and D (Fig. 2a-d). These pits also have the ~~melt~~ enriched LAP layer interleaved between the young and old snow layers, indicated by the sharp increase (or steep slope) between the young and old snow layers. In the two pits where this general outline is not visible (pit E Fig. 2e and pit F Fig. 2f), it can be explained by the fact that pit E extended only to the ~~melt~~ enriched LAP layer (therefore no old snow samples) while pit F had essentially no young snow samples at the time of sampling (therefore pit F starts with the ~~melt~~ enriched LAP layer).

With the data points for young and old snow appearing rather similar in slope between the pits, the homogeneity is emphasized further by comparing the observations with common effective constants for young and old snow (EC_y^* and EC_o^*), respectively. Suitable constants were determined to be close to

50 µg m^{-2} per mm SWE for young snow and 150 µg m^{-2} per mm SWE for old snow (see supplement section 4). The resulting deposition using EC_y^* and EC_o^* are superimposed over the observations in Figure 2a-e as dashed lines for young snow and ~~dash-~~dotted lines for old snow. These lines then represent the constant deposition of EC as function of accumulated melt water in a column according to:

$$ECacc = constant * SWEacc + offset \tag{2}$$

where ECacc is the accumulated EC mass per m^2 and SWEacc is the accumulated melt water in L m^{-2} (or mm), and the 'constant' is the deposition constant. The offsets for young snow are a result of enhanced observed EC concentration in the top layer, which can numerically be compensated for by "artificially" adding a small value to (ΔSWEacc) to each pit (except pit A), which in essence dilute the top layer, but have marginal effect on the overall picture. This meant simply rewriting the linear relation above into:

$$ECacc = constant * (SWEacc + \Delta\text{SWEacc}) \tag{3}$$

The ΔSWEacc amounts were chosen by trial and error to be in multiples of 10 mm for simplicity. The resulting values were ~~10~~, 10, ~~6~~50, 20, and 20 mm for pits ~~A~~ B through E in order to explain the apparent offset. A physical interpretation of these numbers may be the loss of water from the surface layer due to evaporation or sublimation, which enhance $[EC]$ in the top layer. For the old snow layers, snow and EC were numerically removed in the data by subtracting accumulated EC and SWE (including the ~~melt~~ enriched LAP layer, when present) down to the old snow layer. This was done such that the first data point satisfies EC_o^*. Hence, for old snow $[EC]_1 d_{SWE_1}/EC_o^*$=SWE_{acc_1} where the index (1) represents the top layer of old snow.

By applying the offset values and numerically removing the upper snow layers, we compare the data in Fig. 2a-f in two separate figures (Fig. 3a-b), one where young snow are grouped together and one for old snow. In Fig 3a, the observed $ECacc$ is plotted against the $ECacc$ value if EC_y^* is used. In Fig. 3b the observed $ECacc$ is plotted against the $ECacc$ value if EC_o^* is used. Note that for old snow the first data point in the different pits will, by definition, be on the 1:1 line. Nevertheless, the consistency between the pits is striking and the fact that much of the variation in $ECacc$ as function of $SWEacc$ (or depth in the pit) can be explained by EC_y^* and EC_o^* alone is a very interesting finding.

3.2 ~~Melt~~ Enriched LAP layer

On the contrary to the observed similarities in the different pits between young and old snow, the ~~melt~~ enriched LAP layer samples do not display similar trends. Instead of being characterized by a common constant, the $ECacc$ value as function of $SWEacc$ in the ~~melt~~ enriched LAP layer differs by orders of magnitude between the different pit profiles. To explore the ~~melt~~ enriched LAP layers further, we make use of the constant for young snow, EC_y^*. Assuming that this is a characteristic value for precipitation during the winter season, we can estimate the required amount of precipitation ($SWEacc$) that is needed to explain the observed $ECacc$ deposition. These derived precipitation amounts for each pit are presented in Figure 4 as a function of the relative depth from the surface to the bottom of the pit. Using this approach, pit F corresponds to a total equivalent of about 2~~1~~200 mm in precipitation, whereas pits B, E, and D represent ~~35~~2800, ~~43~~200, and ~~51~~4900 mm, respectively. Pits A and C deviate starkly from the others, with 3~~7~~6000 and ~~55~~4000 mm precipitation. Comparing these derived values to other precipitation estimates allows us to provide a temporal perspective required to explain the observed EC in the pits. Other studies have shown that the annual precipitation is very altitude-level dependent in the Himalayas, and based on the altitude of the glaciers alone one would expect less than about 1000 mm in annual precipitation (Anders et al., 2006; Bookhagen and Burbank, 2010). Based on the changes in snow depth, the local precipitation was estimated using the AWS as described in section 2.2. This analysis gave a snow accumulation of about 600 mm SWE in the winter season 2015-2016 and 700 mm in the 2016-2017 winter season at the location of the AWS. Over the season, a fraction of the snow evaporates or sublimates, possibly accounting for a magnitude of mm per day during favorable conditions (Stigter et al., 2018). Further, Mimeau et al. (2019) estimated the sublimation between 12 and 15 % of the total annual precipitation in the Khumbu valley, Nepal. This amount might be missed by this method using daily data. Nonetheless, our two precipitation estimates are below the observed annual precipitation of 976 mm in 2012/2013 at 3950 m altitude, about 250 km to the north-west next to the Chhota Shigri glacier front (Azam et al., 2016). Measured with an automatic precipitation gauge (i.e. capturing all precipitation forms), the authors found that the majority of precipitation was during the winter season, and that the summer monsoon contributed with only 12 % to the annual precipitation. Based on these observation estimates, and the similarities with our Sunderdhunga AWS precipitation patterns, we estimate that about 800 ± 200 mm is a characteristic annual precipitation amount close to where the pits were dug. If the precipitation amounts derived to explain the deposited EC in each pit is divided by 800 mm, the minimum number of years required to explain the EC observed in the pit is acquired. With this approach it is clear that it would require decades of precipitation to explain the EC in the ~~melt~~ enriched LAP layers in pits A and C. This is unrealistic, especially when the lower levels in pit F from the previous year is compared. Even the difference in EC amount between pits B, E, and D compared to F is much greater than can be explained from aggregating the EC accumulated by one year of precipitation in a single melt layer. At the same time, the dry deposition of EC probably accounts for only a few percent of the deposition. With a dry deposition velocity of EC of 0.3 mm/s (Emerson et al.,

2018) and an atmospheric concentration of 0.3 µg m^{-3}, reported at similar altitude at the Nepal Pyramid station during the pre-monsoon (Bonasoni et al., 2010), the dry deposition can be estimated to 2800 µg m^{-2} annually, which several orders of magnitude lower than what is encountered in the enriched LAP layers. Thus, this leads us to propose that EC must have been transported laterally in the surface layer during the melt period in the summer of 2016 and converged in the altitude range where the pits were dug. From Figure 1 it can be seen that the pits were dug in a complex terrain where slopes with increasing gradient are reaching up to the summit towards the southwest.

The data and analysis presented above lead us to propose that the old snow layers observed in pit F from 2015 are the same old snow layers observed for the pits dug in 2016. The EC equivalent precipitation profile of pit F presented in Figure 4 suggests that strong melting had taken place already in summer 2015. Hence, the old snow is composed of snow from at least the season 2013-2014 (or perhaps also earlier seasons). Stratigraphy analysis for pit F presented in Svensson et al. (2018) suggested that the snow deposition represented five seasons. The amount of precipitation represented by the EC deposition (cf. Figure 4) in the old snow is about 2~~1~~200 mm, which suggests that the EC was deposited over several seasons, but less than 5 seasons. Another strong melt took place in 2016, possibly leading to melting all of the snow from the season 2015-2016. In addition, during the melting phase, water and snow particulates could be transported down the slopes from areas of the glacier with steep slopes. Because the steepness of the slope decreases towards the valley, this resulted in a convergence of percolated material from areas above the sampling sites. The young snow is likely part of the 2016-2017 winter season that had started to accumulate before the sampling in October 2016 was commenced. This is confirmed by AWS data that indicates intermittent snow events in October 2016. At the AWS location a seasonal snow cover was in place in December 2016.

3.3 Mineral dust fraction in snow

An initial inspection of the mineral fractional absorption on the filters did not reveal any special common pattern in concentration between the different pits, except for the ~~melt~~ enriched LAP layer samples, which appeared to have higher concentrations than the other samples. In Figure 5, the data is grouped according to the pit stratigraphy classification, and although the absolute range of MD fractions in young snow samples is very large (5 to 71 %), the quartile range is only between 32 to 48 % with a median value of 39 %. The median value for old snow is somewhat larger at 46 %, along with the range and quartiles, which are closer together, from 26 to 70 % and from 43 to 50 %, respectively. The range of values for the ~~melt~~ enriched LAP layer are consistently higher compared to the other two snow types. The median is 78 % with a range and quartiles of 48 to 95 % and 74 to 82 %, respectively. Note that from a total of 95 samples only 16 are from the ~~melt~~ LAP layer. As with EC, MD has the propensity to remain at the snow surface with melting (e.g. Doherty et al., 2013).

Due to the typically heavy loading of material on the filters obtained in the ~~melt~~ enriched LAP layer, those values should be taken with caution, however. Non-linear effects could skew the resulting light absorption fractions towards larger values ~~science~~ since with a very heavy loading (dark filter) the contribution by remaining particles may be over-estimated. This is because the relative contribution by additional light absorbing material decreases as the amount of material increases on very dark filters. In an extreme case, black on black will not add any contribution. The larger range of values in young snow compared to old snow is possibly an effect from the geometric thickness of the sampled slabs, which are in young snow generally thinner than in old snow, and that the density of young snow is typically less than the density of old snow. This results in each of the sampled segments in young snow representing less deposition of both water and LAP and, therefore, presenting a larger variability. Nevertheless, the ensemble of data presents similar median values for both young and old snow. The median of the percentage of the mineral dust absorption f_D value for young and old snow samples together becomes 44 %. The specific absorption by minerals is expected to be orders of magnitude smaller than BC (e.g. Utry et al., 2015), and the same is expected with respect to EC. This suggests that the deposition of minerals in the snow is orders of magnitude larger than EC. If we simply scale our characteristic EC constants (EC_y^* and EC_o^*), with the median of f_D and the ratio between their specific mass absorption coefficients (MAC), according to:

$$\frac{f_D}{(1-f_D)}\frac{MAC_{EC}}{MAC_D}EC_c = D_c \qquad (4)$$

we arrive at a mass concentration for minerals. We use a MAC for BC of 7.5 m^2 g^{-1} (Bond and Bergstrom, 2006). The MAC for the minerals is not known and can vary significantly, but for the sake of this test we use a MAC value representative for the mineral quartz with 0.0023 m^2 g^{-1} (Utry et al., 2015). If we use these values we arrive at a range of 128-384 µg g^{-1} of minerals in the snow. This is in range with previous gravimetric observations from Himalaya (e.g. Thind et al., 2019; Zhang et al., 2018).

3.4 Discussion

Our results indicate that the contribution to light absorption by minerals can be comparable to light absorption by EC in the Sunderdhunga area at about 5 km altitude. This translates into a mass concentration ratio between EC and minerals of more than three orders of magnitude. These large ratios are typically not reported for air samples because much of the deposited minerals are likely from local sources. This supports a hypothesis of a positive climate feedback that results in a reduction of snow cover and the exposure to larger sources of minerals.

For the Tibetan plateau, Zhang et al. (2018) estimated that the retreat of the snow cover could be advanced by more than a week due to LAP in snow. In their estimates, BC accounted for most of this effect and dust advanced the melting by about one day. The BC concentration in snow used in their calculations were about one order of magnitude larger than our derived values form the profiles in the snow pits. This difference can be attributed to the significant contribution of aerosol particle dry deposition in arid regions (Wang et al., 2014), but the range of values presented in their Table 2 reveals a potential problem from sampling surface snow. Post depositional processes (e.g. sublimation/evaporation, hoar formation, snow drift) can alter the concentration at a given location relatively fast, which is less of a problem if a deeper layer of the snow pack is investigated instead of solely the surface snow. Simply taking a larger vertical slab is not sufficient as is evident from the melt layer in the present study. The ~~melt~~ enriched LAP layer in the pits can be studied to characterize the short-term seasonal surface albedo, but the aerosol concentrations cannot be directly related to the deposition. The consistency between pits and different sampling seasons in the integrated deposition profiles above and below the ~~melt~~ enriched LAP layer show the strength in the data collected from snow pits in comparison to snap-shot conditions of surface snow.

4 Conclusions

In this study we aimed at characterizing the observed deposition of EC in the glacier snow in the Sunderdhunga valley and to estimate the contribution from minerals to LAP in the snow. The analysis illustrates that in the sampling area of Durga Kot and Bhanolti glaciers, the deposition of EC in young snow (from current winter season) is characterized by approximately 50 µg m^{-2} mm^{-1} SWE water, which is in the range of other observations. The median fraction of light absorption caused by minerals was about 39 % (Q1=32, Q3=48). In old snow (from previous winter seasons), the deposition was characterized by about 150 µg m^{-2} mm^{-1} SWE water. The reason for this difference can simply be due to a larger deposition in the years before sampling was conducted, or that more water had the chance to leave the snow-pack of older snow. Different from young snow, old snow have had to survive at least one summer season. The median fraction of light absorption was 46 % (Q1=43, Q3=50) by minerals in the old snow layer. Although the variability within each layer is rather large, the obtained lower median fraction for young snow is consistent with the fact that old snow is more exposed to rock surfaces free of snow during the summer season.

Between these two layers of old and young snow, a clearly visible and very dark layer was present. This layer was most likely a result of strong melting that took place in the summers of 2015 and 2016 as discussed in 3.2. However, the high concentration of EC found in this layer cannot simply be explained by a collapse of the snow-pack vertically, and thus it is concluded that lateral transport of LAP (including EC and minerals) took place that resulted in a convergence of material in the altitude range of the snow pits. Different from the other two layers (young and old snow), this melt enriched LAP layer presented large differences with respect to EC content among the different pits. The fraction of light absorption by minerals was the highest of the three layers and was about 80 % (Q1=74, Q3=82).

The profiles of EC and the mineral absorption fraction show good agreement between subsequent years and among different pits. At the same time, the topography in this mountainous region of Himalaya evidently causes great complexity with respect to the distribution of LAP in the snow surface layer during periods of strong melt. Although data is limited in spatial and temporal dimensions our results are useful for large scale radiation impact assessments of EC deposition and minerals. In small scale regional studies, however, the effects of complex topography and spatial variability should be considered separately. Future work should further study the mineral dust and its composition in the area, in order to more accurately elucidate dust role in the snow radiation scheme state in this part of the Himalaya.

Data availability

All data are available upon request.

Author contributions

J. Sv, H.H, E.A., N.D., H.L., participated in the field expedition. S.T., R.H., V.S., M. L., H.L., A.H. handled project administration. Data analysis was performed by J. Sv. and J. St. Funding acquisition: A.H. Supervision M.L. and H.L. J. Sv led the writing of the manuscript with J. St., with input from all other co-authors.

Completing interests.

The authors declare that they have no conflict of interest.

Acknowledgements

This work has been supported by the Academy of Finland project: Absorbing Aerosols and Fate of Indian Glaciers (AAFIG; project number 268004), and the Academy of Finland consortium: "Novel Assessment of Black Carbon in the Eurasian Arctic: From Historical Concentrations and Sources to Future Climate Impacts" (NABCEA project number 296302). J.Svensson acknowledges support from the two Finnish foundations: Maj and Tor Nessling and Oskar Huttunen; as well as the invited scientist grant from the UGA. J. Ström is part of the Bolin Centre for Climate Research, and acknowledges the Swedish Research Council grant 2017-03758. We are thankful for Daniela Tuomala's work with the filter analyzes, as well as the strenuous assistance given by Sherpas and mountain guides during the expeditions to the Sunderdhunga valley.

References

Anders, A. M., Roe, G. H., Hallet, B., Montgomery, D. R., Finnegan, N. J., and Putkonen, J.: Spatial patterns of precipitation and topography in the Himalaya, Geol. Soc. Am. Special Papers, 398, 39–53, 2006.

Azam, M. F., Ramanathan, A. L., Wagnon, P., Vincent, C., Linda, A., Berthier, E., Sharma, P., Mandal, A., Angchuk, T., Singh, V. B., and Pottakkal, J. G.: Meteorological conditions, seasonal and annual mass balances of Chhota Shigri Glacier, western Himalaya, India, Ann. Glaciol., 57, 328–338, doi: 10.3189/2016AoG71A570, 2016.

Bond, T. C., and Bergstrom, R. W.: Light absorption by carbonaceous particles: An investigative review, Aerosol Sci. Tech., 40, 27–67, doi:10.1080/02786820500421521, 2006.

Bonasoni, P., Laj, P., Marinoni, A., Sprenger, M., Angelini, F., Arduini, J., Bonafè, U., Calzolari, F., Colombo, T., Decesari, S., Di Biagio, C., di Sarra, A. G., Evangelisti, F., Duchi, R., Facchini, MC., Fuzzi, S., Gobbi, G. P., Maione, M., Panday, A., Roccato, F., Sellegri, K., Venzac, H., Verza, G. P., Villani, P., Vuillermoz, E., and Cristofanelli, P.: Atmospheric Brown Clouds in the Himalayas: first two years of continuous observations at the Nepal Climate Observatory-Pyramid (5079 m), Atmos. Chem. Phys., 10, 7515–7531, doi:10.5194/acp-10-7515-2010, 2010.

Bookhagen, B. and Burbank, D. W.: Toward a complete Himalayan hydrological budget: Spatiotemporal distribution of snowmelt and rainfall and their impact on river discharge, J. Geophys. Res., 115, F03019, doi:10.1029/2009JF001426, 2010.

Carrico, C. M., Bergin, M. H., Shrestha, A., Dibb, J. E., Gomes, L., and Harris, J.M.: The importance of carbon and mineral dust to seasonal aerosol properties in the Nepal Himalayas, Atmos. Environ., 37, 2811–2824, 2003.

Chow, J. C., Watson, J. G., Crow, S., Lowenthal, D. H., and Merrifield, T.: Comparison of IMPROVE and NIOSH carbon measurements, Aerosol Sci. Tech., 34, 23–34, 2001.

Doherty, S. J., Grenfell, T. C., Forsström, S., Hegg, D. L., Brandt, R. E., and Warren, S. G.: Observed vertical redistribution of black carbon and other insoluble light-absorbing particles in melting snow, J. Geophys. Res.-Atmos., 118, 5553–5569, doi.org/10.1002/jgrd.50235, 2013.

Emerson, E. W., Katich, J. M., Schwarz, J. P., McMeeking, G. R., and Farmer, D. K.: Direct Measurements of Dry and Wet Deposition of Black Carbon Over a Grassland, J. Geophys. Res.-Atmos., 123, 12277–212290, doi.org/10.1029/2018JD028954, 2018.

Flanner, M. G., Zender, C. S., Randerson, J. T., and Rasch, P. J.: Present-day climate forcing and response from black carbon in snow, J. Geophys. Res.-Atmos., 112, D11202, doi.org/10.1029/2006JD008003, 2007.

Gertler, C. G., Puppala, S. P., Panday, A., Stumm, D., and Shea, J.: Black carbon and the Himalayan cryosphere: A review, Atmos. Environ., 125, 404–417, doi.org/10.1016/j.atmosenv.2015.08.078, 2016.

He, C., Flanner, M. G., Chen, F., Barlage, M., Liou, K.-N., Kang, S., Ming, J., and Qian, Y.: Black carbon-induced snow albedo reduction over the Tibetan Plateau: uncertainties from snow grain shape and aerosol–snow mixing state based on an updated SNICAR model, Atmos. Chem. Phys., 18, 11507–11527, doi.org/10.5194/acp-18-11507-2018, 2018.

Helfricht, K., Hartl, L., Koch, R., Marty, C., and Olefs, M.: Obtaining sub-daily new snow density from automated measurements in high mountain regions, Hydrol. Earth Syst. Sci., 22, 2655–2668, doi.org/10.5194/hess-22-2655-2018, 2018.

Hooda, R.K., Kivekäs, N., O'Connor, E.J., Collaud Coen, M., Pietikäinen, J.P., Vakkari, V., Backman, J., Henriksson, S.V., Asmi, E., Komppula, M., Korhonen, H., Hyvärinen, A. P., Lihavainen, H.: Driving factors of aerosol properties over the foothills of central Himalayas based on 8.5 Years continuous measurements, J. Geophys. Res. Atmos., 123, doi.org/10.1029/2018JD029744, 421-13,442, 2018.

Jacobi, H.-W., Lim, S., Ménégoz, M., Ginot, P., Laj, P., Bonasoni, P., Stocchi, P., Marinoni, A., and Arnaud, Y.: Black carbon in snow in the upper Himalayan Khumbu Valley, Nepal: observations and modeling of the impact on snow albedo, melting, and radiative forcing, The Cryosphere, 9, 1685–1699, doi.org/10.5194/tc-9-1685-2015, 2015.

Jacobson, M. Z.: Climate response of fossil fuel and biofuel soot, accounting for soot's feedback to snow and sea ice albedo and emissivity, J. Geophys. Res.-Atmos., 109, D21201, doi.org/10.1029/2004jd004945, 2004.

Kaspari, S., Painter, T. H., Gysel, M., Skiles, S. M., and Schwikowski, M.: Seasonal and elevational variations of black carbon and dust in snow and ice in the Solu-Khumbu, Nepal and estimated radiative forcings, Atmos. Chem. Phys., 14, 8089–8103, doi.org/10.5194/acp-14-8089-2014, 2014.

Krecl, P., Ström, J., and Johansson, C.: Carbon content of atmospheric aerosols in a residential area during the wood combustion season in Sweden, Atmos. Environ., 41, 6974–6985, doi.org/10.1016/j.atmosenv.2007.06.025, 2007.

Kopacz, M., Mauzerall, D. L., Wang, J., Leibensperger, E. M., Henze, D. K., and Singh, K.: Origin and radiative forcing of black carbon transported to the Himalayas and Tibetan Plateau, Atmos. Chem. Phys., 11, 2837–2852, doi.org/10.5194/acp-11-2837-2011, 2011.

Li, C., Bosch, C., Kang, S., Andersson, A., Chen, P., Zhang, Q.,Cong, Z., Chen, B., Qin, D., and Gustafsson, O.: Sources of blackcarbon to the Himalayan-Tibetan Plateau glaciers, Nature Com-mun., 7, 12574, doi.org/10.1038/ncomms12574, 2016.

Ménégoz, M., Krinner, G., Balkanski, Y., Boucher, O., Cozic, A., Lim, S., Ginot, P., Laj, P., Gallée, H., Wagnon, P., Marinoni, A., and Jacobi, H. W.: Snow cover sensitivity to black carbon deposition in the Himalayas: from atmospheric and ice core measurements to regional climate simulations, Atmos. Chem. Phys., 14, 4237–4249, doi.org/10.5194/acp-14-4237-2014, 2014.

Mimeau, L., Esteves, M., Zin, I., Jacobi, H.-W., Brun, F., Wagnon, P., Koirala, D., and Arnaud, Y.: Quantification of different flow components in a high-altitude glacierized catchment (Dudh Koshi, Himalaya): some cryospheric-related issues, Hydrol. Earth Syst. Sci., 23, 3969–3996, doi.org/10.5194/hess-23-3969-2019, 2019.

Ming, J., Du, Z., Xiao, C., Xu, X., and Zhang, D.: Darkening of the mid-Himalaya glaciers since 2000 and the potential causes, Environ. Res. Lett., 7, 014021, doi:10.1088/1748-25 9326/7/1/014021, 2012.

Ming, J., Xiao, C., Du, Z., and Yang, X.: An Overview of Black Carbon Deposition in High Asia Glaciers and its Impacts on Radiation Balance, Adv. Water Resour., 55, 80–87, 2013.

Petzold, A., Ogren, J. A., Fiebig, M., Laj, P., Li, S.-M., Baltensperger, U., Holzer-Popp, T., Kinne, S., Pappalardo, G., Sugimoto, N., Wehrli, C., Wiedensohler, A., and Zhang, X.-Y.: Recommendations for reporting "black carbon" measurements, Atmos. Chem. Phys., 13, 8365–8379, doi.org/10.5194/acp-13-8365-2013, 2013.

Raatikainen, T., Hyvärinen, A. P., Hatakka, J., Panwar, T. S., Hooda, R. K., Sharma, V. P., and Lihavainen, H.: The effect of boundary layer dynamics on aerosol properties at the Indo-Gangetic plains and at the foothills of the Himalayas, Atmos. Environ., 89, 548– 555, doi.org/10.1016/j.atmosenv.2014.02.058, 2014.

Ramanathan, V. and Carmichael, G.: Global and regional climate changes due to black carbon, Nat. Geosci., 1, 221–227, doi.org/10.1038/ngeo156, 2008.

Santra, S., Verma, S., Fujita, K., Chakraborty, I., Boucher, O., Takemura, T., Burkhart, J. F., Matt, F., and Sharma, M.: Simulations of black carbon (BC) aerosol impact over Hindu Kush Himalayan sites: validation, sources, and implications on glacier runoff, Atmos. Chem. Phys., 19, 2441–2460, doi.org/10.5194/acp-19-2441-2019, 2019.

Sarangi, C., Qian, Y., Rittger, K., Bormann, K. J., Liu, Y., Wang, H., Wan, H., Lin, G., and Painter, T. H.: Impact of light-absorbing particles on snow albedo darkening and associated radiative forcing over high-mountain Asia: high-resolution WRF-Chem modeling and new satellite observations, Atmos. Chem. Phys., 19, 7105–7128, doi.org/10.5194/acp-19-7105-2019, 2019.

Skiles, S. M., Flanner, M., Cook, J. M., Dumont, M., and Painter, T. H.: Radiative forcing by light-absorbing particles in snow, Nat. Clim. Change, 8, 964–971, doi.org/10.1038/s41558-018-0296-5, 2018.

Stigter, E. E., Maxime Litt, A., Jakob Steiner, F., Pleun Bonekamp, N. J., Joseph Shea, M., and Immerzeel, W. W.: The importance of snow sublimation on a himalayan glacier. Front. Earth Sci. 6:108, doi: 10.3389/feart.2018.00108, 2018.

Svensson, J., Ström, J., Kivekäs, N., Dkhar, N. B., Tayal, S., Sharma, V. P., Jutila, A., Backman, J., Virkkula, A., Ruppel, M., Hyvärinen, A., Kontu, A., Hannula, H.-R., Leppäranta, M., Hooda, R. K., Korhola, A., Asmi, E., and Lihavainen, H.: Light-absorption of dust and elemental carbon in snow in the Indian Himalayas and the Finnish Arctic, Atmos. Meas. Tech., 11, 1403–1416, doi.org/10.5194/amt-11-1403-2018, 2018.

Thind, P. S., Chandel, K. K., Sharma, S. K., Mandal, T. K., and John, S.: Light-absorbing impurities in snow of the Indian Western Himalayas: impact on snow albedo, radiative forcing, and enhanced melting, Environ. Sci. Pollut. Res. Int., 26, 7566–7578, doi.org/10.1007/s11356-019-04183-5, 2019.

Utry, N., Ajtai, T., Pintér, M., Tombácz, E., Illés, E., Bozóki, Z., and Szabó, G.: Mass-specific optical absorption coefficients and imaginary part of the complex refractive indices of mineral dust components measured by a multi-wavelength photoacoustic spectrometer, Atmos. Meas. Tech., 8, 401–410, doi.org/10.5194/amt-8-401-2015, 2015.

Wang, Z. W., Gallet, J. C., Pedersen, C. A., Zhang, X. S., Ström, J., and Ci, Z. J.: Elemental carbon in snow at Changbai Mountain, northeastern China: concentrations, scavenging ratios, and dry deposition velocities, Atmos. Chem. Phys., 14, 629–640, doi.org/10.5194/acp-14-629-2014, 2014.

Warren, S. and Wiscombe, W.: A model for the spectral albedo of snow II. Snow containing atmospheric aerosols, J. Atmos. Sci., 37, 2734–2745, 1980.

Xu, B., Cao, J., Hansen, J., Yao, T., Joswiak, D.R., Wang, N., Wu, G.,Wang, M., Zhao, H., Yang,W., Liu, X., and He, J.: Black soot and the survival of Tibetan glaciers, P. Natl. Acad. Sci. USA, 106, 22114–22118, doi.org/10.1073/pnas.0910444106, 2009.

Zhang, R., Wang, H., Qian, Y., Rasch, P. J., Easter, R. C., Ma, P.-L., Singh, B., Huang, J., and Fu, Q.: Quantifying sources, transport, deposition, and radiative forcing of black carbon over the Himalayas and Tibetan Plateau, Atmos. Chem. Phys., 15, 6205–6223, doi.org/10.5194/acp-15-6205-2015, 2015.

Zhang, Y., Kang, S., Sprenger, M., Cong, Z., Gao, T., Li, C., Tao, S., Li, X., Zhong, X., Xu, M., Meng, W., Neupane, B., Qin, X., and Sillanpää, M.: Black carbon and mineral dust in snow cover on the Tibetan Plateau, The Cryosphere, 12, 413–431, doi.org/10.5194/tc-12-413-2018, 2018.

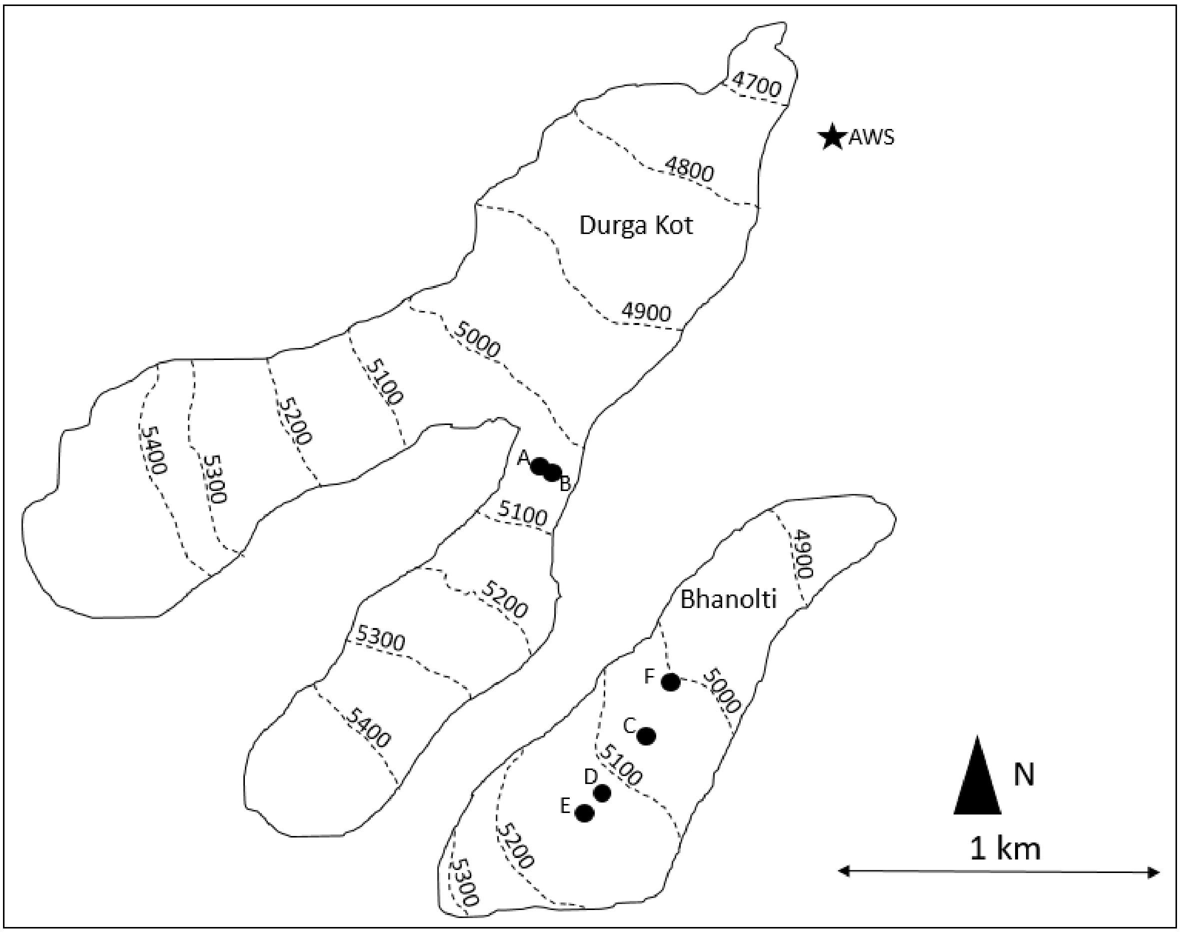
Table 1. Snow pit details from Sunderdhunga valley. Durga Kot glacier snow pits are A-B, while C-F are from Bhanolti glacier.

Snow pit ID and elevation (m a.s.l)	Depth interval (cm)	Snow density (g cm^{-3})		Water equivalent (mm m^{-2})	TC analyzed (µg L^{-1})	EC (µg L^{-1})		EC deposition (µg m^{-2})	*f*D (%)
		Measured	Assumed			Analyzed	Reconstructed		
A, 5055	0-3	0.38		11.4	1130	-	120	1364	24.6
	3-6	0.38		11.4	238	18	-	207	29.2
	6-9	0.35		10.5	477	47	-	495	40.4
	9-12	0.37		11.1	30300	-	3125	34688	-
	12-15	0.39		11.7	1307404	-	134685	1575819	76.1
	15-20		0.50	25.0	68177	-	7034	175855	55.1
	20-25		0.50	25.0	1398	278	-	6945	47.9
	25-30		0.50	25.0	1549	147	-	3684	49.8
	30-35		0.50	25.0	1769	271	-	6787	41.9
	35-40		0.50	25.0	1466	251	-	6273	46.5
	40-45		0.50	25.0	883	141	-	3528	44.6
	45-50		0.50	25.0	751	142	-	3553	43.1
	50-60		0.50	50.0	1090	171	-	8544	51.5
	60-70		0.50	50.0	763	88	-	4412	45.9
B, 5055	0-3	0.40		12.0	1542	95	-	1143	38.3
	3-6	0.40		12.0	693	30	-	364	27.5
	6-9	0.39		11.6	31710	-	3291	38015	77.8
	9-12	0.33		9.9	69667	-	7210	71378	75.0
	12-15	0.33		9.9	3498	-	374	3699	50.6
	15-19		0.50	20.0	-	-	267	5348	49.9
	19-29		0.50	50.0	1534	246	-	12319	49.8
	29-39		0.50	50.0	1295	190	-	9480	46.2
	39-49		0.50	50.0	1517	248	-	12407	52.1

	49-59		0.50	50.0	1753	182	-	9100	40.2
	59-69		0.50	50.0	733	103	-	5156	41.2
	69-79		0.50	50.0	730	102	-	5121	44.9
C, 5068	0-3	0.40		12.0	2386	-	249	2983	47.6
	3-6	0.39		11.7	590	45	-	523	31.6
	6-11	0.39		19.5	372	34	-	658	59.4
	11-16	0.42		21.0	799	93	-	1959	54.8
	16-21	0.46		23.0	1074	141	-	3240	58.0
	21-26		0.50	25.0	1047065	-	107865	2696629	-
	26-31		0.50	25.0	4480	370	-	9257	62.0
	31-36		0.50	25.0	684	80	-	1988	58.9
	36-41		0.50	25.0	906	150	-	3746	43.6
	41-46		0.50	25.0	658	126	-	3159	44.2
	46-56		0.50	50.0	863	137	-	6871	43.5
	56-66		0.50	50.0	1191	156	-	7803	45.7
	66-76		0.50	50.0	832	144	-	7222	44.9
	76-86		0.50	50.0	802	94	-	4709	45.9
	86-96		0.50	50.0	416	51	-	2543	42.6
	96-106		0.50	50.0	609	78	-	3913	45.3
	106-116		0.50	50.0	692	76	-	3821	50.0
	116-126		0.50	50.0	500	46	-	2322	57.9
	126-136		0.50	50.0	1265	108	-	5386	59.0
D, 5125	0-3	0.39		11.7	1135	127	-	1487	35.2
	3-6	0.39		11.7	1012	91	-	1068	34.0
	6-9	0.37		11.1	449	30	-	337	42.6
	9-12	0.37		11.1	810	41	-	450	47.3
	12-15	0.37		11.0	1089	84	-	916	48.5
	15-18	0.37		11.0	357	32	-	353	38.6
	18-21	0.36		10.8	918	59	-	637	38.5
	21-24	0.42		12.6	274		36	448	70.8
	24-27	0.42		12.6	322	23	-	293	57.2

	27-30	0.36		10.8	443	28	-	297	36.3
	33-36	0.36		10.8	2393		253	2734	95.1
	36-39	0.45		13.5	1714		186	2506	77.6
	39-42	0.45		13.5	6806		710	9591	77.1
	42-44		0.50	10.0	177424		18313	183125	-
	44-49		0.50	20.0	9733		1025	20504	60.1
	49-54		0.50	25.0	5708	665	-	16635	59.5
	54-59		0.50	25.0	1743	232	-	5798	69.6
	59-69		0.50	50.0	901	129	-	6459	46.1
E, 5143	0-3	0.33		9.9	992	128	-	1268	35.9
	3-6	0.33		9.9	422	63	-	622	41.4
	6-9	0.37		11.1	891	81	-	903	25.9
	9-12	0.31		9.3	569	41	-	380	43.0
	12-15	0.31		9.3	806	73	-	681	27.7
	15-18	0.29		8.7	750	35	-	302	41.0
	18-21	0.29		8.7	345	22	-	193	55.6
	21-24	0.39		11.7	644		81	943	4.5
	24-27	0.38		11.4	500	50	-	566	27.0
	27-30	0.38		11.4	439	65	-	739	56.7
	30-33	0.40		12.0	395	53	-	635	49.4
	33-36	0.40		12.0	642	26	-	308	27.3
	36-39	0.44		13.2	397	33	-	430	38.9
	39-42	0.44		13.2	1250	53	-	705	34.8
	42-45	0.44		13.2	1148	75	-	988	48.4
	45-48	0.45		13.5	828	169	-	2287	81.1
	48-51	0.45		13.5	901	131	-	1775	77.7
	51-54	0.45		13.5	617	58	-	786	85.8
	54-55	0.45		4.5	-	-	4694	21125	85.8
	55-60		0.50	25.0	69606	-	7198	179946	85.8
F, 5008	0-3		0.35	10.5	4075	66		690	77.9
	3-6		0.40	12.0	4821	273		3271	60.8

6-10	0.40	16.0	17686		1828	29242	69.5
10-15	0.50	25.0	3555	233		5830	60.3
15-20	0.50	25.0	859	111		2786	33.1
20-30	0.50	50.0	1324	141		7036	49.2
30-40	0.50	50.0	807	106		5278	39.6
40-50	0.50	50.0	890	98		4907	36.7
50-60	0.50	50.0	2825	270		13484	49.5
60-70	0.50	50.0	1228	179		8965	39.9
70-80	0.50	50.0	696	93		4650	36.1
80-90	0.50	50.0	483	73		3640	35.8
90-100	0.50	50.0	1190	144		7190	43.9
100-110	0.50	50.0	652	79		3965	29.5
110-120	0.50	50.0	554	57		2846	25.7


4700
AWS
4800
Durga Kot
4900
5000
5100
5200
5400
5300
A
B
5100
5200
5300
5400
Bhanolti
4900
F
5000
C
5100
D
E
5200
5300
N
1 km

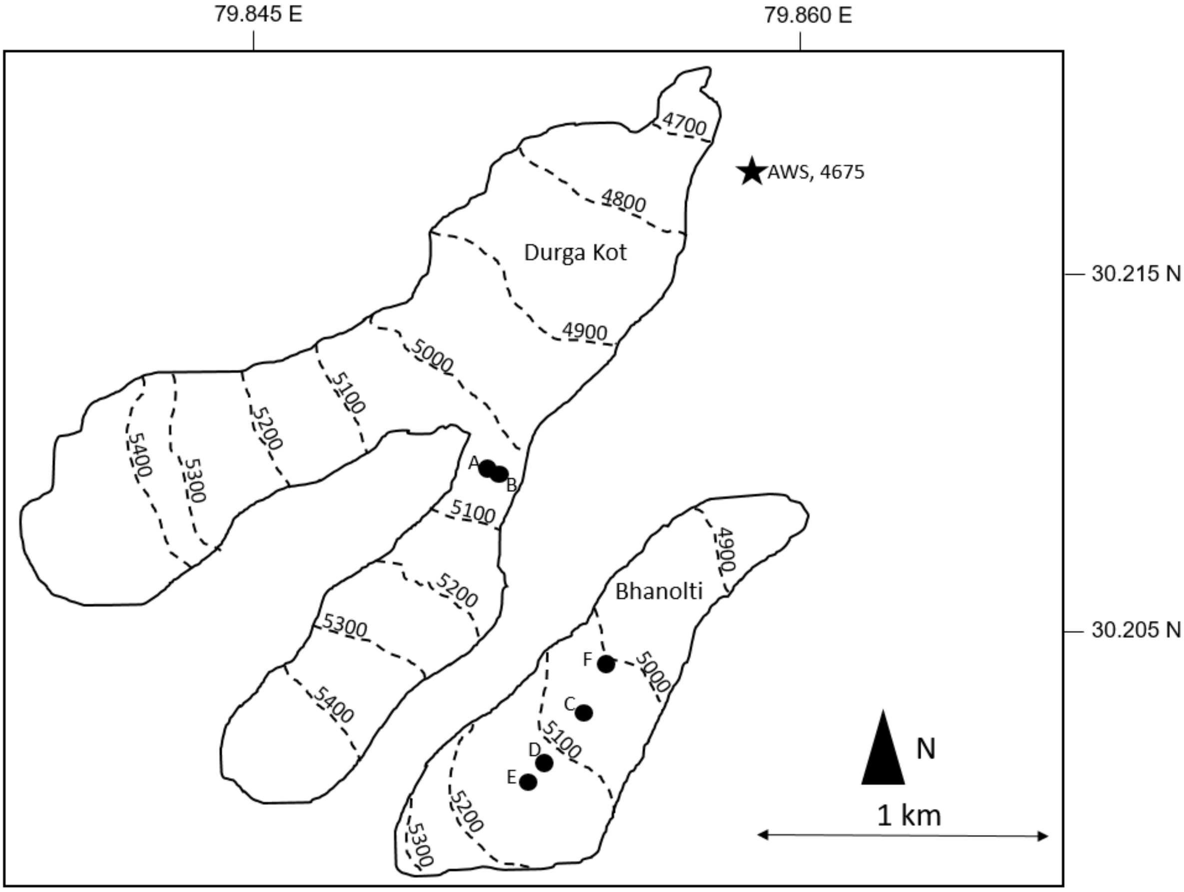
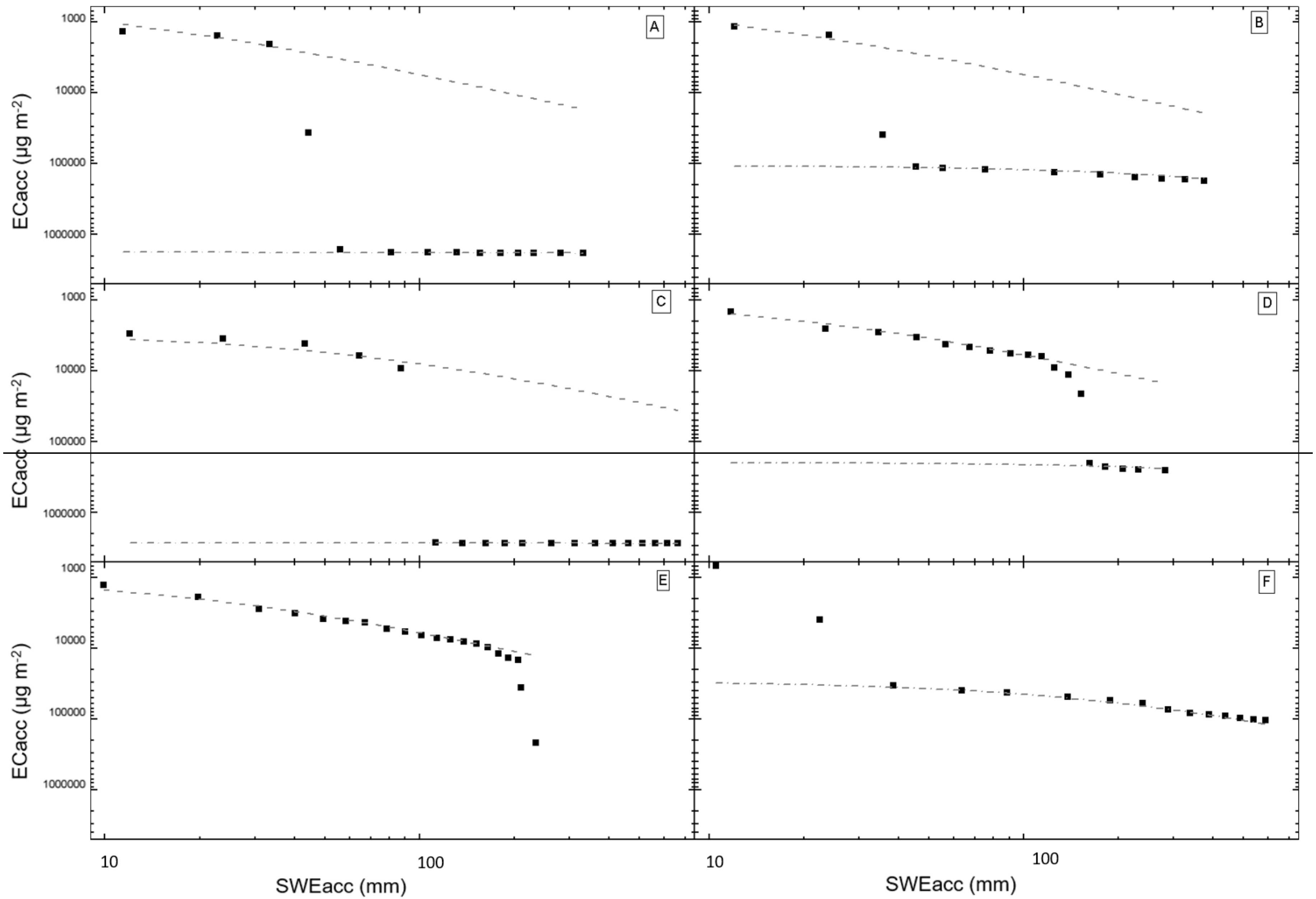
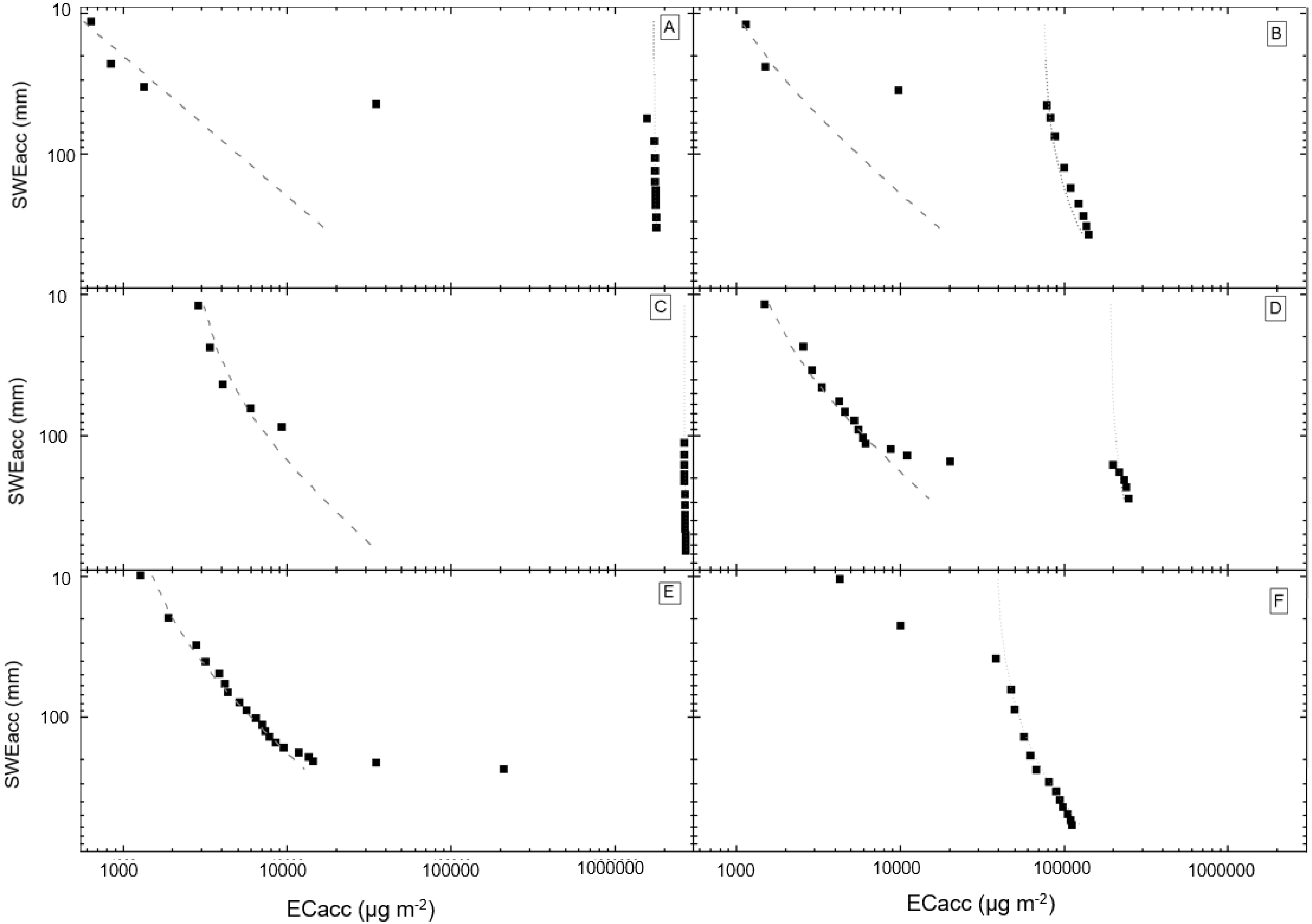

79.845 E
79.860 E
4700
AWS, 4675
4800
Durga Kot
30.215 N
4900
5000
5100
5200
5400
5300
A
B
5100
4900
Bhanolti
5200
30.205 N
5300
F
5000
5400
C
5100
D
N
E
5200
1 km
5300

Figure 1. Map of glaciers with the location of the snow pits (black dots) and AWS indicated with a star. Dashed lines on the glacier refer to iso-lines .


A
B
C
D
E
F
1000
10000
100000
1000000
ECacc (µg m-2)
10
100
SWEacc (mm)

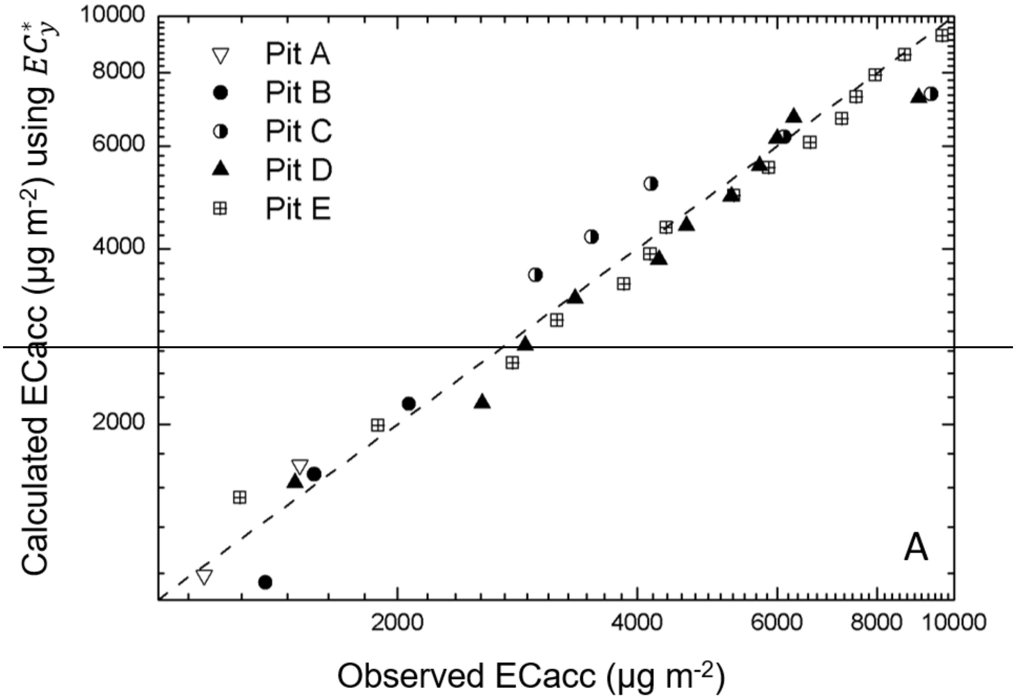
570

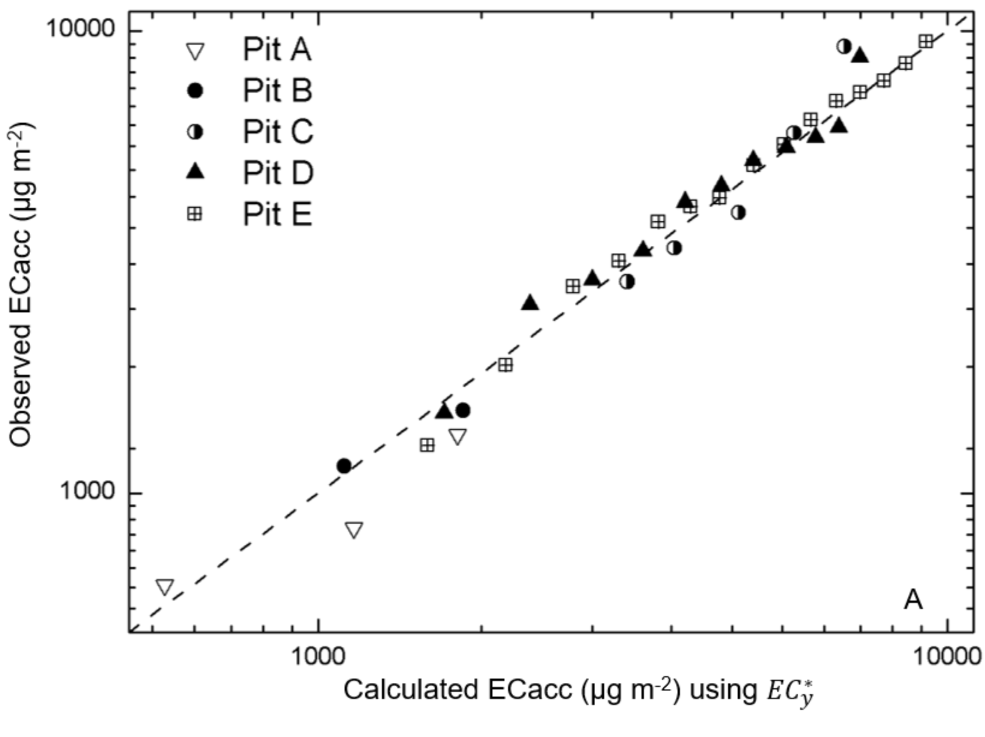
571

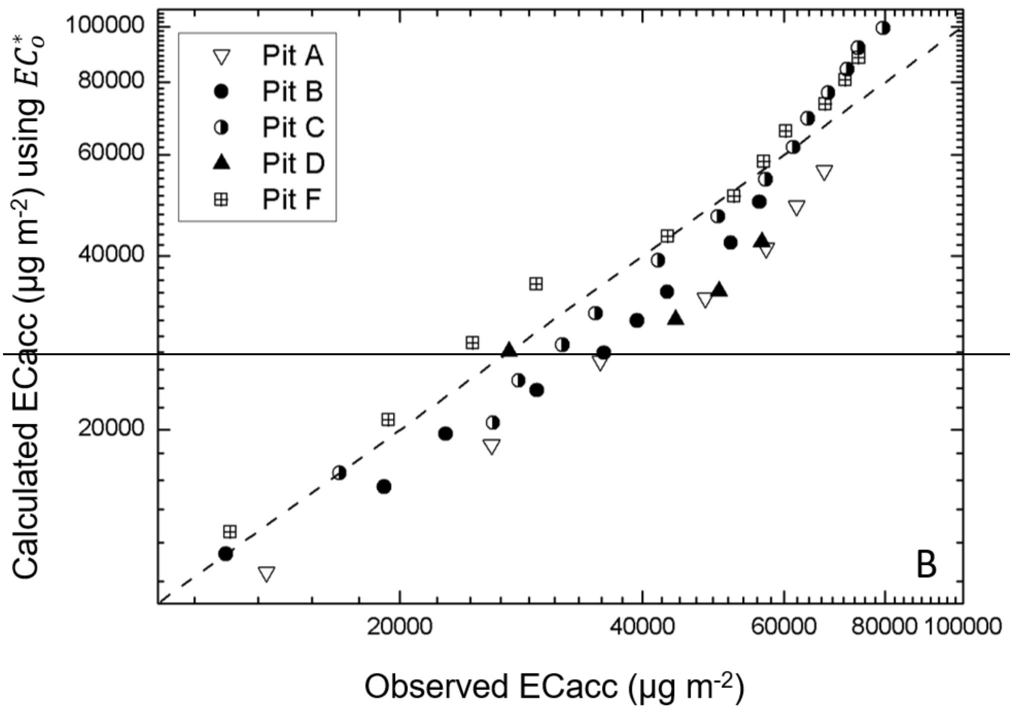


A
B
C
D
E
F
SWEacc (mm)
SWEacc (mm)
SWEacc (mm)
10
100
1000
10000
100000
1000000
ECacc (µg m-2)
ECacc (µg m-2)

Figure 2. The cumulative $\widetilde{EC}_i$(ECacc) from top to bottom in the snow pits as function of accumulated d_{SWEi} expressed as SWEacc (mm): (a) Pit A, (b) Pit B, (c) Pit C, (d) Pit D, (e) Pit E, (f) Pit F. The upper dashed line represents a constant deposition EC_y^* and the lower dashed-dotted line represents a constant deposition EC_o^*. In pit E there were no snow samples classified as old snow, hence there is no n EC_o^* line, while in in pit F there were no young snow samples, therefore no EC_y^* line.


Pit A
Pit B
Pit C
Pit D
Pit E
A
10000
8000
6000
4000
2000
2000
4000
6000
8000
10000
Calculated ECacc (μg m-2) using EC_y^*
Observed ECacc (μg m-2)


Pit A
Pit B
Pit C
Pit D
Pit E
A
10000
1000
1000
10000
Observed ECacc (μg m-2)
Calculated ECacc (μg m-2) using EC_y^*



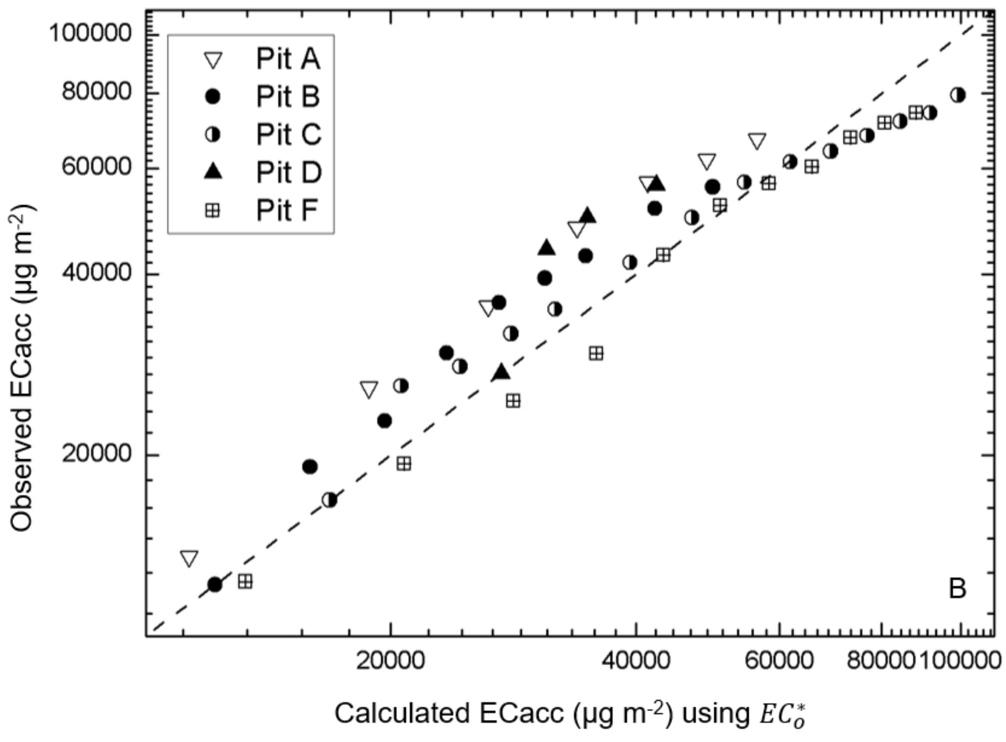


Figure 3. Observed and the calculated deposition using the constant deposition EC_y^* for young (a) and EC_o^* for old (b) snow samples. Dashed lines indicate a 1:1 slope.

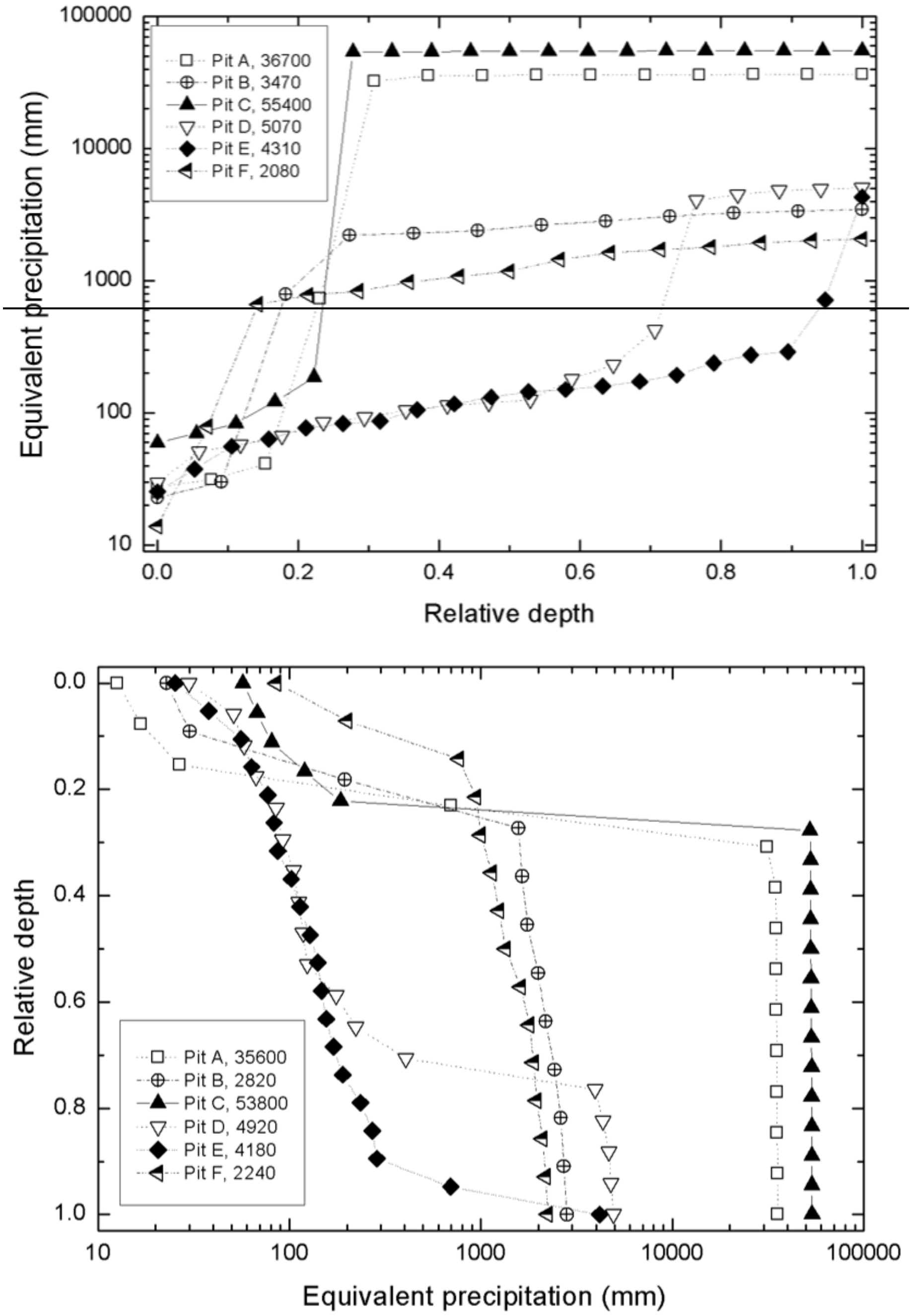
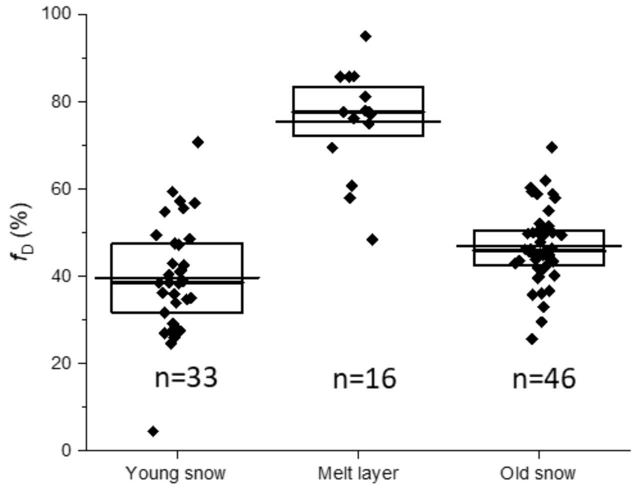


Figure 4. Equivalent precipitation for each pit based on a constant deposition EC_y^* in fresh snow as function of the relative depth of the pit from top to bottom.



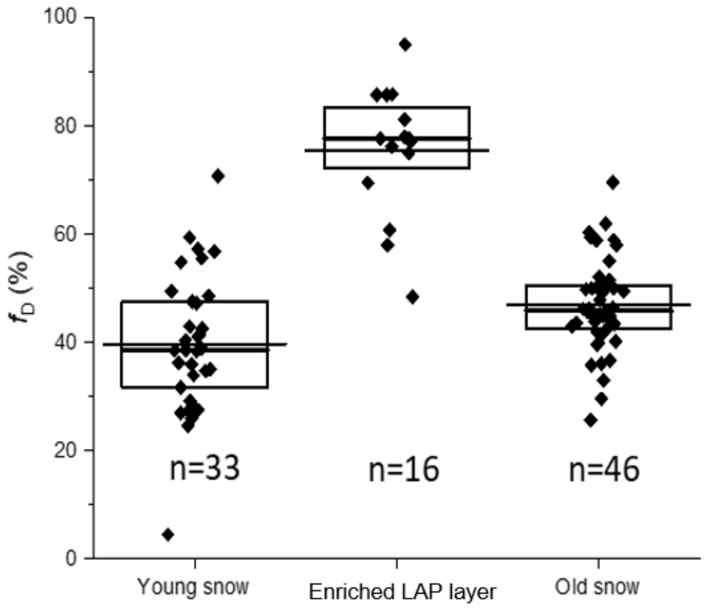


Figure 5. Fractional dust absorption remaining after burning the filters during OC/EC analysis. The diamonds are individual values for each filter and the thin extended line represents the arithmetic average. The box and thicker line represent the quartile range and median, respectively. The number of samples are indicated in the figure as (n).