

1 **Technical note: Measurement of chemically-resolved volume**
2 **equivalent diameter and effective density of particles by AAC-**
3 **SPAMS**

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17 **Abstract**

18 Size and effective density (ρ_e) are important properties of aerosol particles and are
19 related to their influences on human health and the global climate. The volume
20 equivalent diameter (D_{ve}) is an intrinsic property that is used to evaluate particle size.
21 Three definitions of ρ_e are generally used to characterize the physical property of a
22 particle as an alternative to particle density, in which only the ρ_e^{II} , defined as the ratio
23 of particle density (ρ_p) to a dynamic shape factor (χ), has the characteristic of being
24 independent of particle size. However, it is still challenging to simultaneously
25 characterize the D_{ve} and ρ_e^{II} of aspherical particles. Here, we present a novel system
26 that classifies particles with their aerodynamic diameter (D_a) by aerodynamic aerosol
27 classifiers (AAC) and determines their vacuum aerodynamic diameter (D_{va}) by single
28 particle aerosol mass spectrometry (SPAMS) to achieve a measurement of D_{ve} and ρ_e^{II} .
29 The reliability of the AAC-SPAMS system for accurately obtaining D_{ve} and ρ_e^{II} is
30 verified based on the results that the deviation between the measured and theoretical
31 values is less than 6% for the size-resolved spherical polystyrene latex (PSL). The
32 AAC-SPAMS system is applied to characterize the D_{ve} and ρ_e of $(\text{NH}_4)_2\text{SO}_4$ and NaNO_3
33 particles, suggesting that these particles are aspherical and their ρ_e are independent of
34 particle size. Finally, the AAC-SPAMS system is deployed in a field measurement,
35 showing that it is a powerful technique to characterize the chemically-resolved D_{ve} and
36 ρ_e^{II} of particles in real time.

37 1. Introduction

38 Size and particle density (ρ_p) are critical parameters of aerosol particles in
39 quantifying the impact of aerosols on air quality, human health and global climate
40 change (Buseck and Posfai, 1999; Poschl, 2005; Pitz et al., 2003). Effective density (ρ_e)
41 has been adopted to characterize the physical property of a particle as an alternative to
42 ρ_p , since ρ_p for aspherical aerosol particles is hardly measured (Sumlin et al., 2018;
43 Katrib et al., 2005). Size and ρ_e govern the transport properties of a particle both in the
44 atmosphere and in the human respiratory system (Seinfeld and Pandis, 1998; Liu and
45 Daum, 2008) and directly and/or indirectly influence the potential of the particle to
46 absorb or reflect solar radiation (Tang, 1997; Zhao et al., 2019; Liu and Daum, 2008).
47 ρ_e can also provide information concerning particle morphology (Yon et al., 2015) and
48 serve as a tracer for atmospheric processing (Guo et al., 2014; Yin et al., 2015; Liu et
49 al., 2015). However, the quantitative relationship between aerosol properties, namely,
50 size and ρ_e , and their effects on air quality, human health and global climate change is
51 not yet well understood, which is partly because important aerosol properties cannot be
52 measured by current techniques.

53 **Size.** Size is a fundamental property of particles, which can be parameterized by the
54 physical quantity of volume equivalent diameter (D_{ve}). Defined as the diameter of a
55 spherical particle with the same volume as the particle (DeCarlo et al., 2004), D_{ve} is an
56 intrinsic physical quantity that can be used to evaluate the actual size of the particle.
57 However, to date, atmospheric science usually describes particle size by other diameter
58 definitions, such as the electric mobility diameter (D_m), aerodynamic equivalent

59 diameter (D_a) and vacuum aerodynamic equivalent diameter (D_{va}), whose relationships
 60 with D_{ve} are shown in Eqs. (1)-(3), respectively:

$$61 \quad \frac{D_m}{C_c(D_m)} = \frac{D_{ve}}{C_c(D_{ve})} \chi_t, \quad (1)$$

$$62 \quad D_a = D_{ve} \sqrt{\frac{\rho_p C_c(D_{ve})}{\chi_t \rho_0 C_c(D_a)}}, \quad (2)$$

$$63 \quad D_{va} = \frac{\rho_p D_{ve}}{\rho_0 \chi_v}, \quad (3)$$

64 where $C_c(D)$ is the Cunningham slip correction factor, χ_t and χ_v represent the aerosol
 65 dynamic shape factor (χ) in the transition regime and in the free-molecule regime,
 66 respectively, and ρ_0 represents the unit density of 1.0 g/cm³. From these definitions, it
 67 can be seen that D_m , D_a , and D_{va} are originally derived from D_{ve} , but in actuality, they
 68 do not reflect the actual size of the aspherical particle. Meanwhile, D_{ve} of aspherical
 69 particles cannot be easily obtained, which limits its application in the scientific
 70 community.

71 **Effective density.** At present, three definitions of ρ_e are introduced in atmospheric
 72 science (DeCarlo et al., 2004): the first definition (ρ_e^I) is the ratio of the measured
 73 particle mass (m_p) to the particle volume (V) calculated assuming a spherical particle
 74 with a diameter equal to the measured D_m ; the second definition (ρ_e^{II}) is the ratio of ρ_p
 75 to χ (Hand and Kreidenweis, 2002); and the third definition (ρ_e^{III}) is the ratio of D_m and
 76 D_{va} , all of which are expressed in Eqs. (4)-(6), respectively.

$$77 \quad \rho_e^I = \frac{6m_p}{\pi D_m^3} \quad (4)$$

$$78 \quad \rho_e^{II} = \frac{\rho_p}{\chi} \quad (5)$$

$$79 \quad \rho_e^{III} = \frac{D_{va}}{D_m} \rho_0 \quad (6)$$

80 The definitions of ρ_e^I and ρ_e^{III} can be derived into the final forms, as shown in the Eqs.(7)

81 and (8), respectively.

$$82 \quad \rho_e^I = \frac{\rho}{\chi t^3} \cdot \left(\frac{C_c(D_{ve})}{C_c(D_m)} \right)^3 \quad (7)$$

$$83 \quad \rho_e^{III} = \rho \cdot \frac{C_c(D_{ve})}{\chi^2 \cdot C_c(D_m)} \quad (8)$$

84 The Eq. (7) is derived from combining the Eq. (1) with Eq. (4), in which m_p is equal to
85 $1/6 \rho \cdot D_{ve}^3$. The detailed derivation of Eq. (8) was presented in Schneider et al. (2006). A
86 variety of methods are developed to characterize ρ_e^I and ρ_e^{III} , among which the more
87 advanced methods are to achieve the measurement of the chemically-resolved effective
88 density. Combining a single particle soot photometer (SP2) with a (volatility) tandem
89 differential mobility analyser ((VT)DMA) can measure the ρ_e^I of particles mixed with
90 soot (Zhang et al., 2016b; Wu et al., 2019; Han et al., 2019). The measurement of
91 chemically-resolved ρ_e^{III} can be achieved by coupling a DMA with an on-line aerosol
92 mass spectrometer such as Single Particle Laser Ablation Time-of-Flight Mass
93 Spectrometer (SPLAT) (Zelenyuk et al., 2005; Zelenyuk et al., 2006; Alexander et al.,
94 2016), an aerosol mass spectrometer (AMS) (Dinar et al., 2006; Schneider et al., 2006;
95 Kiselev et al., 2010), an aerosol time-of-flight mass spectrometer (ATOFMS) (Spencer
96 and Prather, 2006; Spencer et al., 2007), and single-particle aerosol mass spectrometry
97 (SPAMS) (Zhang et al., 2016a; Zhai et al., 2017). However, the ρ_e^I and ρ_e^{III} are
98 demonstrated to have the inherent characteristics of decreasing with increasing particle
99 size, which will be presented in a separate publication. Therefore, it will introduce
100 systemic error when assessing the particle impacts on visibility, human health and
101 climate change from the physical quantities in ρ_e^I and ρ_e^{III} . In contrast, ρ_e^{II} is independent
102 of particle size. For example, for soot particles with χ of 2.5 and ρ_p of 1.80 g/cm³, the

103 calculated ρ_e^I , ρ_e^{II} , and ρ_e^{III} are 0.43, 0.72, and 0.45 g/cm³ at D_m of 40 nm, and 0.22,
104 0.72, and 0.36 g/cm³ at D_m of 550 nm, respectively. The big gap between the three
105 definitions of effective density suggests that they should be carefully treated when
106 characterizing the particles. However, the ρ_e^{II} has not been widely applied in
107 atmospheric sciences because of the lack of measurement techniques. Previous
108 literatures tried to retrieve the ρ_e^{II} and the real part in the refractive index (n) through a
109 fitting procedure that compares the measured light-scattering intensity of particles
110 (R_{meas}) to the theoretical values ($R_{theory,test}$) calculated by a series of n and ρ_e^{II} values
111 (Moffet and Prather, 2005; Moffet et al., 2008; Zhang et al., 2016a). Moffet and Prather
112 (2005) successfully obtained ρ_e^{II} for spherical particles by single particle mass
113 spectrometry. However, subject to the accuracy of Mie theory for the aspherical
114 particles, dry NaCl and calcium-rich dust particles were failed to fit the $R_{theory,test}$ well
115 to R_{meas} (Moffet et al., 2008). Similarly, Zhang et al. (2016a) failed to simultaneously
116 retrieve ρ_e^{II} and n for (NH₄)₂SO₄ and NaNO₃ particles. To our best knowledge, there is
117 no appropriate technique to achieve the measurement of ρ_e^{II} for aspherical particles.

118 The aim of the present work is to develop a method to simultaneously obtain D_{ve} and
119 ρ_e^{II} for aspherical particles. For simplicity, the symbol ρ_e in the following text refers to
120 the definition of ρ_e^{II} . The established system of an aerodynamic aerosol classifier
121 (AAC)-SPAMS is capable of characterizing the D_a and D_{va} of particles, which can be
122 applied to theoretically derive D_{ve} and ρ_e . To verify the reliability of the AAC-SPAMS
123 system, we apply it to measure the D_{ve} and ρ_e of the spherical particles of polystyrene
124 latex (PSL). The results are in good agreement with the theoretical values. Finally, the

125 AAC-SPAMS system is applied to measure the D_{ve} and ρ_e for $(\text{NH}_4)_2\text{SO}_4$ and NaNO_3
126 particles and for the chemically-resolved atmospheric particles.

127

128 **2. Experimental section**

129 **2.1 Measurement system**

130 Figure 1 shows a schematic diagram of the AAC-SPAMS system. The particles are
131 first dried by a diffusion drying tube (TSI 9302, USA), classified by AAC (Cambustion
132 Ltd., UK) based on the aerodynamic diameters D_a , and then transported into SPAMS in
133 which the D_{va} and the mass spectra of individual particles are obtained. The working
134 principle of the AAC is described in detail elsewhere (Tavakoli and Olfert, 2013). AAC
135 consists of two coaxial cylinders that rotate at the same rotational speed. Polydisperse
136 particles enter into the space between the cylinders (i.e., classification column) and
137 experience a centrifugal force that causes them to move toward the outer cylinder. The
138 particles to be classified can leave the classification column with the particle-free sheath
139 flow and finally exit the AAC with the sample flow. Thus, the D_a values of classified
140 particles can be derived from their relationship with their relaxation time (τ), as shown
141 in Eq. (9):

$$142 \quad \tau = \frac{C_C(D_a) \cdot \rho_0 \cdot D_a^2}{18\mu} \quad (9)$$

143 where μ is the gas dynamic viscosity. Particles with large relaxation times impact and
144 adhere to the outer cylinder, while particles with small relaxation times exit the
145 classifier with the exhaust flow. The exhaust flow from the AAC was about 0.3 lpm,
146 and the Size Resolution Parameter (Rs) of the AAC was set as 40.

147 Detailed information about the operation of SPAMS (Hexin Analytical Instrument
148 Co., Ltd., China) is given elsewhere (Li et al., 2011). Briefly, the particles are
149 introduced into the vacuum system through a 0.1 mm critical orifice and are gradually
150 collimated into a beam in the aerodynamic lens. Two continuous diode Nd:YAG laser
151 beams (532 nm) are used to aerodynamically size the particles, which are subsequently
152 desorbed/ionized by a pulsed laser (266 nm) that is triggered based on the velocity of a
153 specific particle. The generated positive and negative ions are recorded with the
154 corresponding particle size. The D_{va} of the particle is related to the transit time between
155 the two laser beams (532 nm) in SPAMS, which can be obtained by using a calibration
156 curve generated from the measured transit times of a PSL series with predefined sizes
157 (nominal diameters).

158

159 **2.2 Laboratory experiments**

160 Dried spherical PSL (Nanosphere Size Standards, Duke Scientific Corp., Palo Alto)
161 ($\rho_p = 1.055 \text{ g/cm}^3$ and $\chi = 1.0$) with D_{ve} values of $203 \pm 5 \text{ nm}$, $310 \pm 6 \text{ nm}$, $510 \pm 5 \text{ nm}$,
162 and $740 \pm 6 \text{ nm}$ were used in the AAC-SPAMS system, and the D_{ve} was verified by
163 Scanning Mobility Particles Sizer (Model 3938, TSI Inc., USA). The PSL particles were
164 first classified by AAC, and then their D_{va} values were obtained by SPAMS. ACC-
165 SPAMS was also applied to the particles of $(\text{NH}_4)_2\text{SO}_4$ ($\rho_p = 1.77 \text{ g/cm}^3$) and NaNO_3
166 ($\rho_p = 2.26 \text{ g/cm}^3$) with D_a values of 250.0 nm, 350.0 nm, 450.0 nm and 550.0 nm.
167 Besides, to present the measurement uncertainty of the AAC, the D_a values of these
168 PSL particles were measured to be 212.8 ± 0.2 , 324.7 ± 0.4 , 529.9 ± 0.4 , and $767.5 \pm$

169 0.4 by the system of AAC- condensation particle counter (CPC), respectively. It shows
170 that the AAC has the deviations of 1.1%, 1.3%, 0.8%, and 0.7% for determining the D_a
171 values of the particles.

172 **2.3 Ambient sampling**

173 For field observations, the AAC-SPAMS system was deployed in Science and
174 Technology Enterprise Accelerator A2 Block, Guangzhou, China, to characterize the
175 D_{ve} , ρ_e and chemical compositions of aerosol particles. The sampling inlet was hung 2.5
176 meters from the third floor (~12 m above ground level). Ambient aerosol particles were
177 introduced into the AAC through a 5 m long conductive silicone tube with an inner
178 diameter of 6 mm and a PM_{2.5} cyclone inlet. The sampling flow from the PM_{2.5} cyclone
179 inlet was 3 lpm, and the residence time in the conductive silicone tube was
180 approximately 5 seconds. Particles with the D_a of 250.0, 350.0, 450.0, and 550.0 nm
181 were classified by the AAC. The sampling time for the particles of each D_a was
182 approximately 10 minutes. From July 6th to 8th, 2019, approximately 129,869 ionized
183 particles were obtained from nine rounds of measurement. The sampling details are
184 shown in Table S1. The number of ionized particles with the D_a of 250.0, 350.0, 450.0,
185 and 550.0 nm is 35,609, 38,374, 31,910, and 23,976, respectively. The sampled
186 ~100,000 particles are first classified by using an adaptive resonance theory neural
187 network (ART-2a) (Song et al., 1999) with a vigilance factor of 0.75, a learning rate of
188 0.05 and 20 iterations.

189

190 **2.4 Theoretical derivation of D_{ve} and ρ_e from D_a and D_{va}**

191 In this study, the calculations of D_{ve} and ρ_e for unknown particles are theoretically
 192 derived from D_a and D_{va} . Combining Eqs. (2) and (3), we obtain the following Eq. (10):

$$193 \quad C_c(D_a) \frac{D_a^2}{D_{va}} = D_{ve} C_c(D_{ve}) \frac{\chi_v}{\chi_t} \quad (10)$$

194 Based on the approximation between χ_v and χ_t ($\chi_v \approx \chi_t = \chi_a$) (DeCarlo et al., 2004), Eq.
 195 (10) becomes Eq. (11):

$$196 \quad C_c(D_a) \frac{D_a^2}{D_{va}} = D_{ve} C_c(D_{ve}) \quad (11)$$

197 The Cunningham Slip Correction Factor is calculated by Eq. (12) (Peng and Bi, 2020):

$$198 \quad C_c(D) = 1 + \frac{\lambda}{D} \left(A + B \cdot \exp\left(\frac{C \cdot D}{\lambda}\right) \right), \quad (12)$$

199 where λ is the mean free path of the gas molecules, and A , B and C are empirically
 200 determined constants specific to the analysis system. The values of A , B and C are 2.33,
 201 0.966, and -0.498 provided by the manual of the AAC. Substituting Eq. (12) into Eq.
 202 (11) obtains the Eq. (13).

$$203 \quad \frac{D_a^2}{D_{va}} + \frac{D_a \cdot \lambda}{D_{va}} \left(A + B \cdot \exp\left(\frac{C \cdot D_a}{\lambda}\right) \right) = D_{ve} + \lambda \left(A + B \cdot \exp\left(\frac{C \cdot D_{ve}}{\lambda}\right) \right) \quad (13)$$

204 If the D_a and D_{va} of an unknown particle can be measured, its D_{ve} could be calculated
 205 according to Eq. (13). Finally, the ρ_e value of the particles is calculated by the D_{va} and
 206 D_{ve} values according to Eq. (14), which is obtained by combining Eq.(3) and Eq.(5):

$$207 \quad \rho_e = \frac{\rho_p}{\chi_a} = \frac{D_{va}}{\rho_0 \cdot D_{ve}} \quad (14)$$

208 Thus, we can obtain both the D_{ve} and ρ_e values of unknown particles based on the D_a
 209 and D_{va} values. Because the AAC and SPAMS instruments have the ability to determine
 210 D_a and D_{va} , the AAC-SPAMS system developed in this study can be used to obtain the
 211 D_{ve} and ρ_e values for unknown particles.

212

213 3. Results and discussion

214 3.1 Verification of the AAC-SPAMS system to obtain D_{ve} and ρ_e

215 The D_{va} distribution of PSL particles with predefined D_{ve} values after screened by the
216 AAC is shown in Figure S1. We used Gaussian fitting to obtain the peak D_{va} for each
217 size PSL with an R-squared fitting coefficient (R^2) over 0.98. Each fitting has a full
218 width at half maximum (FWHM) of 6.6%, 4.4%, 2.3% and 2.2%, and the corresponding
219 peaks are 215.8 nm, 319.0 nm, 532.1 nm and 803.5 nm, respectively. Substituting the
220 D_a and D_{va} values of PSL into Eq. (11), the measured D_{ve} ($D_{ve,me}$) of PSL from AAC-
221 SPAMS system is 203.6 nm, 309.7 nm, 511.6 nm and 737.2 nm, respectively (Figure
222 2a). Thus, the deviations between the theoretical D_{ve} ($D_{ve,th}$) and $D_{ve,me}$ values are 0.3%,
223 -0.1%, 0.3% and -0.4%, respectively. On the other hand, the measured ρ_e ($\rho_{e,me}$) values
224 of the particles, calculated from the D_{va} and $D_{ve,me}$ values with Eq. (14), are 1.1 g/cm³,
225 1.0 g/cm³, 1.0 g/cm³, and 1.1 g/cm³, respectively (Figure 2b). Comparing to the
226 theoretical ρ_e ($\rho_{e,th}$) (i.e. 1.055 g/cm³ of PSL particles), the deviations of $\rho_{e,me}$ are
227 determined to be 4.3%, -5.2%, -5.2%, and 4.3%, respectively. That is, the deviations of
228 $D_{ve,me}$ and $\rho_{e,me}$ obtained by the AAC-SPAMS system are within 1% and 6%,
229 respectively.

230

231 3.2 Application of the AAC-SPAMS system for obtaining D_{ve} and ρ_e of $(\text{NH}_4)_2\text{SO}_4$ 232 and NaNO_3

233 Figure S2 shows the D_{va} distributions of $(\text{NH}_4)_2\text{SO}_4$ and NaNO_3 particles with D_a
234 values of 250.0, 350.0, 450.0, and 550.0 nm screened by the AAC. The D_{va} peaks are

235 obtained by Gaussian fitting, with R^2 values over 0.93 and FWHM values ranging from
236 7.6% to 10.6%. The $(\text{NH}_4)_2\text{SO}_4$ particles have D_{va} values of 300.0, 418.0, 551.1, and
237 695.1 nm (Figure S2), which correspond to particles possessing $D_{ve,me}$ values of 177.3,
238 254.4, 331.8, and 409.3 nm, respectively, according to Eq. (11). Substituting the values
239 of D_{va} and $D_{ve,me}$ into Eq. (12), the $\rho_{e,me}$ values are 1.7, 1.6, 1.6, and 1.7 g/cm^3 (Figure
240 3a), respectively. Similarly, the selected NaNO_3 particles are determined to have D_{va}
241 values of 321.0, 454.9, 599.8, and 755.3 nm (Figure S2), corresponding to $D_{ve,me}$ values
242 of 150.1, 218.2, 287.0, and 355.9 nm, respectively. The $\rho_{e,me}$ values of the NaNO_3
243 particles are 2.2, 2.0, 2.0, and 2.1 g/cm^3 (Figure 3b), respectively. Figure 3 also shows
244 that the $\rho_{e,me}$ values of the NaNO_3 and $(\text{NH}_4)_2\text{SO}_4$ particles with different D_a deviate
245 from their average values with the maximum of 5.9 % and 4.8%, respectively, which
246 are identical with the deviation for the $\rho_{e,me}$ of PSL particles. These deviations may be
247 derived from the calibration of particle D_{va} from the SPAMS. While the R-square of
248 size calibration curve is 0.999, the curve of exponential function is found to slightly
249 deviate from the data points measured by SPAMS. For example, size calibration
250 function produces the deviation of -4.4% and 3.1% from the data points of 310 and 740
251 nm, respectively.

252 Taking the systematic error into account, the slight difference of the $\rho_{e,me}$ values for
253 the four sizes suggests that the ρ_e of $(\text{NH}_4)_2\text{SO}_4$ and NaNO_3 particles is independent of
254 particle size from 250.0 nm to 550.0 nm. It is determined by the definition of effective
255 density used in this study, which keeps constant as long as the χ_a of the particles does
256 not change with particle size for pure compound. The average $\rho_{e,me}$ values of $(\text{NH}_4)_2\text{SO}_4$

257 and NaNO₃ particles are calculated to be 1.7 ± 0.1 and 2.1 ± 0.1 g/cm³, which are lower
258 than the ρ_p of (NH₄)₂SO₄ (1.77 g/cm³) and NaNO₃ (2.27 g/cm³). This is partly caused
259 by the χ_a , which can be used to parameterize the morphology. According to Eq. (14),
260 the χ_a with different D_a are calculated to be 1.04, 1.11, 1.11, and 1.04 for (NH₄)₂SO₄
261 particles and to be 1.03, 1.14, 1.14, and 1.08 for NaNO₃ particles. Thus, the average χ_a
262 values of the (NH₄)₂SO₄ and NaNO₃ particles are determined to be 1.07 ± 0.04 and 1.10
263 ± 0.05 , respectively, indicating that these particles are aspherical.

264 The asphericity of (NH₄)₂SO₄ determined by AAC-SPAMS system is consistent with
265 the previous studies reporting that the χ_a of (NH₄)₂SO₄ were larger than the value of
266 1.03 (Zelenyuk et al., 2006; Beranek et al., 2012; Zhang et al., 2016a). However,
267 previous studies found that the NaNO₃ particles had different morphology. Zhang et al.
268 (2016a) observed that NaNO₃ had the χ_a of 1.09-1.13, while Hoffman et al. (2004) found
269 that NaNO₃ particle had a round droplet-like shape even at 15% RH, supported by the
270 consistence between the measured value of “anhydrous” droplet density and the
271 calculated value of “anhydrous” solution droplet (Zelenyuk et al., 2005). Eclectically,
272 Tang and Munkelwitz (1994) studied that most of the NaNO₃ particles crystallized
273 between 20% and 30% RH but some persisted down to 10% RH to keep solution
274 droplets. Notably, the spherical NaNO₃ particles at low RH observed by Hoffman et al.
275 (2004) were dried in the sticky carbon tape which might affect the phase transition of
276 droplet-like NaNO₃ particles. In this study, most NaNO₃ particles might crystallize
277 because the RH of the aerosol flow carrying the NaNO₃ particles was reduced to below
278 20% through the diffusion drying tube. The asphericity of the crystallized NaNO₃

279 particles is supported by their FWHM values of the D_{va} distributions, which are
280 consistent with that of aspherical $(\text{NH}_4)_2\text{SO}_4$ (Figures S1 and S2).

281

282 **3.3 Application of the AAC-SPAMS system for measuring the chemically-resolved**

283 **D_{ve} and ρ_e**

284 SPAMS can obtain information on the chemical composition of individual particles,
285 implying that the AAC-SPAMS system has the ability to simultaneously characterize
286 D_{ve} , ρ_e and the chemical compositions of particles in real time. It is worth noting that
287 the freshly emitted soot particles exhibit the largest χ (~ 2.5) in the actual atmosphere
288 (Peng et al., 2016). It meets the upper limit for the approximation between the χ_t and χ_v
289 (DeCarlo et al., 2004).

290 As an example, the AAC-SPAMS system was deployed in the field to obtain the
291 chemically-resolved D_{ve} and ρ_e values for unknown aerosol particles. The sampled
292 $\sim 100,000$ particles are classified into eight major particle types with distinct chemical
293 composition: K-rich, EC-S, K-Na, Amine, EC-N-S, OC-N-S and OC-EC-N-S and
294 Metal-rich, representing 97% of the detected particle population. Details of the
295 chemical composition and number fraction of the eight types of particles are presented
296 in the Figure S3 and Figure S4, respectively, which are discussed in the Supporting
297 Information.

298 We used Gaussian fitting to obtain the D_{va} peaks for each particle type with D_a values
299 of 250.0 nm, 350.0 nm, 450.0 nm, and 550.0 nm. Then, we calculated the D_{ve} values of
300 the atmospheric particles with Eq. (11). Table 1 presents the average D_{ve} values of the

301 eight particle types, for which the standard deviation is calculated based on nine
302 samples. The average D_{ve} at D_a values of 250.0 nm, 350.0 nm, 450.0 nm, and 550.0 nm
303 shows wide ranges: from 188.5 nm to 200.8 nm, 271.9 nm to 295.7 nm, 342.5 nm to
304 428.9 nm, and 397.3 nm to 570.9 nm, respectively, which are caused by the different
305 chemical composition. The result indicates that particles with significantly different D_{ve}
306 might possess the same D_a . Furthermore, the large standard deviation of D_{ve} , such as
307 21.9 nm for K-Na at 250.0 nm, 32.3 nm for OC-EC-N-S at 350.0 nm, and 44.3 nm for
308 OC-N-S at 450.0 nm, indicates that the D_{ve} of particles is remarkably different even for
309 particles with the same type and same D_a .

310 According to D_{ve} and D_{va} , we calculated the ρ_e of each particle type by Eq. (12).
311 Figure 4 shows the variations of the ρ_e with D_{ve} for nine particle samples. For pure
312 compounds, such as $(\text{NH}_4)_2\text{SO}_4$ and NaNO_3 particle, ρ_e theoretically does not change
313 with particle size. However, the sampled particles have experienced complex
314 atmospheric processes. Therefore, ρ_e has a very wide distribution for each type of
315 particles with a similar D_{ve} . Specifically, the ρ_e of K-Na increases with D_{ve} , while the ρ_e
316 of OC-N-S and OC-EC-N-S decreases with D_{ve} , which may be influenced by the
317 particle shape or the material density. Additionally, the average ρ_e of each type of
318 particle is in the order from small to large: $1.2 \pm 0.2 \text{ g/cm}^3$ for OC-EC-N-S, 1.3 ± 0.2
319 g/cm^3 for OC-N-S, $1.4 \pm 0.1 \text{ g/cm}^3$ for K-rich, $1.4 \pm 0.1 \text{ g/cm}^3$ for Amine, 1.5 ± 0.1
320 g/cm^3 for EC-N-S, $1.5 \pm 0.1 \text{ g/cm}^3$ for EC-S, $1.6 \pm 0.1 \text{ g/cm}^3$ for K-Na and 1.6 ± 0.1
321 g/cm^3 for Metal-rich. It is reasonable to find that the average ρ_e of internally mixed
322 particles distributes in the range of their material densities (ρ_m). For instance, the OC-

323 EC-N-S, OC-N-S, K-rich, and Amine particles, mainly comprised of internally mixed
324 sulfate and organics, have the average ρ_e between that of sulfate with ρ_m of 1.77 g/cm^3
325 and organic aerosols with ρ_m of 1.2 g/cm^3 (Cross et al., 2007).

326

327 **4. Conclusion**

328 We develop an AAC-SPAMS system to first achieve the measurement of the D_{ve} and
329 ρ_e (defined as the ratio of ρ_p to χ) of the aspherical particles through characterizing their
330 D_a and D_{va} . The reliability of the AAC-SPAMS system is verified by accurately
331 measuring the D_{ve} and ρ_e of PSL. Applying the AAC-SPAMS system to determine the
332 D_{ve} and ρ_e of $(\text{NH}_4)_2\text{SO}_4$ and NaNO_3 particles shows that these particles are aspherical
333 and their ρ_e are independent of particle size. Coupled with the ability of SPAMS to
334 characterize the chemical composition of individual particles, the AAC-SPAMS system
335 was demonstrated to be capable of characterizing the D_{ve} , ρ_e (ρ_p/χ) and chemical
336 compositions of atmospheric particles simultaneously, showing the potential
337 application of this system in field observations. The approach achieves the
338 measurement of chemically-resolved D_{ve} and ρ_e (ρ_p/χ), and provides the possibility to
339 determine their quantitative relationship with other particle properties, which would be
340 benefit for further reduction of the uncertainty associated with the effects of particles
341 on air quality, human health and radiative forcing.

342

343 **Data availability.** Data in this study is available at [https://github.com/longer1217/All-](https://github.com/longer1217/All-figures-data)
344 figures-data.

345

346 **Author contributions.** The idea for the study was conceived by LP and GHZ. All
347 experiments were performed by LP with the assistance of LL. LP wrote the paper which
348 was reviewed by GHZ and XHB. All co-authors discussed the results and commented
349 on the manuscript.

350

351 **Competing interests.** The authors declare they have no conflict of interest.

352

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361 **References**

362 Alexander, J. M., Bell, D. M., Imre, D., Kleiber, P. D., Grassian, V. H., and Zelenyuk,
363 A.: Measurement of size-dependent dynamic shape factors of quartz particles in
364 two flow regimes, *Aerosol Sci. and Technol.*, 50, 870-879,
365 <https://doi.org/10.1080/02786826.2016.1200006>, 2016.

366 Beranek, J., Imre, D., and Zelenyuk, A.: Real-time shape-based particle separation and

367 detailed in situ particle shape characterization, *Anal. Chem.*, 84, 1459-1465,
368 <https://doi.org/10.1021/ac202235z>, 2012.

369 Buseck, P. R., and Posfai, M.: Airborne minerals and related aerosol particles: effects
370 on climate and the environment, *P. Natl. Acad. Sci. USA*, 96, 3372-3379,
371 <https://doi.org/10.1073/pnas.96.7.3372>, 1999.

372 Cross, E. S., Slowik, J. G., Davidovits, P., Allan, J. D., Worsnop, D. R., Jayne, J. T.,
373 Lewis, D. K., Canagaratna, M., and Onasch, T. B.: Laboratory and ambient particle
374 density determinations using light scattering in conjunction with aerosol mass
375 spectrometry, *Aerosol Sci. and Technol.*, 41, 343-359,
376 <https://doi.org/10.1080/02786820701199736>, 2007.

377 DeCarlo, P. F., Slowik, J. G., Worsnop, D. R., Davidovits, P., and Jimenez, J. L.: Particle
378 morphology and density characterization by combined mobility and aerodynamic
379 diameter measurements. Part 1: Theory, *Aerosol Sci. and Technol.*, 38, 1185-1205,
380 <https://doi.org/10.1080/027868290903907>, 2004.

381 Dinar, E., Mentel, T. F., and Rudich, Y.: The density of humic acids and humic like
382 substances (HULIS) from fresh and aged wood burning and pollution aerosol
383 particles, *Atmos. Chem. Phys.*, 6, 5213-5224, [https://doi.org/10.5194/acp-6-5213-](https://doi.org/10.5194/acp-6-5213-2006)
384 2006, 2006.

385 Guo, S., Hu, M., Zamora, M. L., Peng, J. F., Shang, D. J., Zheng, J., Du, Z. F., Wu, Z.,
386 Shao, M., Zeng, L. M., Molina, M. J., and Zhang, R. Y.: Elucidating severe urban
387 haze formation in China, *P. Natl. Acad. Sci. USA*, 111, 17373-17378, 2014.

388 Han, C., Li, S. M., Liu, P., and Lee, P.: Size dependence of the physical characteristics

389 of particles containing refractory black carbon in diesel vehicle exhaust, Environ.
390 Sci. & Technol., 53, 137-145, <https://doi.org/10.1021/acs.est.8b04603>, 2019.

391 Hand, J. L., and Kreidenweis, S. M.: A new method for retrieving particle refractive
392 index and effective density from aerosol size distribution data, Aerosol Sci. and
393 Technol., 36, 1012-1026, <https://doi.org/10.1080/02786820290092276>, 2002.

394 Hoffman, R. C., Laskin, A., and Finlayson-Pitts, B. J.: Sodium nitrate particles:
395 physical and chemical properties during hydration and dehydration, and
396 implications for aged sea salt aerosols, J. Aerosol Sci., 35, 869-887, 2004.

397 Katrib, Y., Martin, S. T., Rudich, Y., Davidovits, P., Jayne, J. T., and Worsnop, D. R.:
398 Density changes of aerosol particles as a result of chemical reaction, Atmos. Chem.
399 Phys., 5, 275-291, <https://doi.org/10.5194/acp-5-275-2005>, 2005.

400 Kiselev, A., Wennrich, C., Stratmann, F., Wex, H., Henning, S., Mentel, T. F., Kiendler-
401 Scharr, A., Schneider, J., Walter, S., and Lieberwirth, I.: Morphological
402 characterization of soot aerosol particles during LACIS Experiment in November
403 (LExNo), J. Geophys. Res.-Atmos., 115, Artn D11204.
404 <https://doi.org/10.1029/2009jd012635>, 2010.

405 Li, L., Huang, Z. X., Dong, J. G., Li, M., Gao, W., Nian, H. Q., Fu, Z., Zhang, G. H.,
406 Bi, X. H., Cheng, P., and Zhou, Z.: Real time bipolar time-of-flight mass
407 spectrometer for analyzing single aerosol particles, Int. J. Mass Spectrom., 303,
408 118-124, <https://doi.org/10.1016/j.ijms.2011.01.017>, 2011.

409 Liu, Y., and Daum, P. H.: Relationship of refractive index to mass density and self-
410 consistency of mixing rules for multicomponent mixtures like ambient aerosols, J.

411 Aerosol Sci., 39, 974-986, <https://doi.org/10.1016/j.jaerosci.2008.06.006>, 2008.

412 Liu, Z., Hu, B., Ji, D., Wang, Y., Wang, M., and Wang, Y.: Diurnal and seasonal
413 variation of the PM_{2.5} apparent particle density in Beijing, China, *Atmos.*
414 *Environ.*, 120, 328-338, <https://doi.org/10.1016/j.atmosenv.2015.09.005>, 2015.

415 Moffet, R. C., and Prather, K. A.: Extending ATOFMS measurements to include
416 refractive index and density, *Anal. Chem.* 77, 6535-6541,
417 <https://doi.org/10.1021/ac0503097>, 2005.

418 Moffet, R. C., Qin, X., Rebotier, T., Furutani, H., and Prather, K. A.: Chemically
419 segregated optical and microphysical properties of ambient aerosols measured in
420 a single-particle mass spectrometer, *J. Geophys. Res.-Atmos.*, 113,
421 <https://doi.org/10.1029/2007jd009393>, 2008.

422 Peng, J. F., Hu, M., Guo, S., Du, Z. F., Zheng, J., Shang, D. J., Zamora, M., Zeng, L.
423 M., Shao, M., Wu, Y. S., Zheng, J., Wang, Y., Glen, C., Collins, D., Molina, M.,
424 and Zhang, R. Y.: Markedly enhanced absorption and direct radiative forcing of
425 black carbon under polluted urban environments, *P. Natl. Acad. Sci. USA*, 252,
426 2016.

427 Peng, L., and Bi, X.: Comment on “Retrieval of atmospheric fine particulate density
428 based on merging particle size distribution measurements: multi-instrument
429 observation and quality control at Shouxian” by Li et al, *J. Geophys. Res.-Atmos.*,
430 125, e2019JD031806, [10.1029/2019JD031806](https://doi.org/10.1029/2019JD031806), 2020.

431 Pitz, M., Cyrys, J., Karg, E., Wiedensohler, A., Wichmann, H. E., and Heinrich, J.:
432 Variability of apparent particle density of an urban aerosol, *Environ. Sci. &*

433 Technol., 37, 4336-4342, <https://doi.org/10.1021/es034322p>, 2003.

434 Poschl, U.: Atmospheric aerosols: Composition, transformation, climate and health
435 effects, *Angew. Chem. Int. Edit.*, 44, 7520-7540,
436 <https://doi.org/10.1002/anie.200501122>, 2005.

437 Schneider, J., Weimer, S., Drewnick, F., Borrmann, S., Helas, G., Gwaze, P., Schmid,
438 O., Andreae, M. O., and Kirchner, U.: Mass spectrometric analysis and
439 aerodynamic properties of various types of combustion-related aerosol particles,
440 *Int. J. Mass Spectrom.*, 258, 37-49, <https://doi.org/10.1016/j.ijms.2006.07.008>,
441 2006.

442 Seinfeld, J. H., and Pandis, S. N.: From air pollution to climate change, 429-443, 1998.

443 Song, X. H., Hopke, P. K., Ferguson, D. P., and Prather, K. A.: Classification of single
444 particles analyzed by ATOFMS using an artificial neural network, ART-2A, *Anal.*
445 *Chem.*, 71, 860-865, <https://doi.org/10.1021/ac9809682>, 1999.

446 Spencer, M. T., and Prather, K. A.: Using ATOFMS to determine OC/EC mass fractions
447 in particles, *Aerosol Sci. and Technol.*, 40, 585-594,
448 <https://doi.org/10.1080/02786820600729138>, 2006.

449 Spencer, M. T., Shields, L. G., and Prather, K. A.: Simultaneous measurement of the
450 effective density and chemical composition of ambient aerosol particles, *Environ.*
451 *Sci. & Technol.*, 41, 1303-1309, <https://doi.org/10.1021/es061425+>, 2007.

452 Sumlin, B. J., Oxford, C. R., Seo, B., Pattison, R. R., Williams, B. J., and Chakrabarty,
453 R. K.: Density and homogeneous internal composition of primary brown carbon
454 *Aerosol, Environ. Sci. & Technol.*, 52, 3982-3989,

455 <https://doi.org/10.1021/acs.est.8b00093>, 2018.

456 Tang, I. N., and Munkelwitz, H. R.: Water activities, densities, and refractive-indexes
457 of aqueous sulfates and sodium-nitrate droplets of atmospheric importance, *J.*
458 *Geophys. Res.-Atmos.*, 99, 18801-18808, 1994.

459 Tang, I. N.: Thermodynamic and optical properties of mixed-salt aerosols of
460 atmospheric importance, *J. Geophys. Res.-Atmos.*, 102, 1883-1893, 1997.

461 Tavakoli, F., and Olfert, J. S.: An instrument for the classification of aerosols by particle
462 relaxation time: theoretical models of the aerodynamic aerosol classifier, *Aerosol*
463 *Sci. and Technol.*, 47, 916-926, <https://doi.org/10.1080/02786826.2013.802761>,
464 2013.

465 Wu, Y. F., Xia, Y. J., Huang, R. J., Deng, Z. Z., Tian, P., Xia, X. G., and Zhang, R. J.: A
466 study of the morphology and effective density of externally mixed black carbon
467 aerosols in ambient air using a size-resolved single-particle soot photometer (SP2),
468 *Atmos. Meas. Tech.*, 12, 4347-4359, 2019.

469 Yin, Z., Ye, X. N., Jiang, S. Q., Tao, Y., Shi, Y., Yang, X., and Chen, J. M.: Size-resolved
470 effective density of urban aerosols in Shanghai, *Atmos. Environ.*, 100, 133-140,
471 <https://doi.org/10.1016/j.atmosenv.2014.10.055>, 2015.

472 Yon, J., Bescond, A., and Ouf, F. X.: A simple semi-empirical model for effective
473 density measurements of fractal aggregates, *J. Aerosol Sci.*, 87, 28-37,
474 <https://doi.org/10.1016/j.jaerosci.2015.05.003>, 2015.

475 Zelenyuk, A., Cai, Y., Chieffo, L., and Imre, D.: High precision density measurements
476 of single particles: The density of metastable phases, *Aerosol Sci. and Technol.*,

477 39, 972-986, <https://doi.org/10.1080/02786820500380206>, 2005.

478 Zelenyuk, A., Cai, Y., and Imre, D.: From agglomerates of spheres to irregularly shaped
479 particles: Determination of dynamic shape factors from measurements of mobility
480 and vacuum aerodynamic diameters, *Aerosol Sci. and Technol.*, 40, 197-217,
481 <https://doi.org/10.1080/02786820500529406>, 2006.

482 Zhai, J. H., Lu, X. H., Li, L., Zhang, Q., Zhang, C., Chen, H., Yang, X., and Chen, J.
483 M.: Size-resolved chemical composition, effective density, and optical properties
484 of biomass burning particles, *Atmos. Chem. Phys.*, 17, 7481-7493,
485 <https://doi.org/10.5194/acp-17-7481-2017>, 2017.

486 Zhang, G., Bi, X., Han, B., Qiu, N., Dai, S., Wang, X., Sheng, G., and Fu, J.:
487 Measurement of aerosol effective density by single particle mass spectrometry,
488 *Science China Earth Sciences*, 59, 320-327, [https://doi.org/10.1007/s11430-015-](https://doi.org/10.1007/s11430-015-5146-y)
489 5146-y, 2016a.

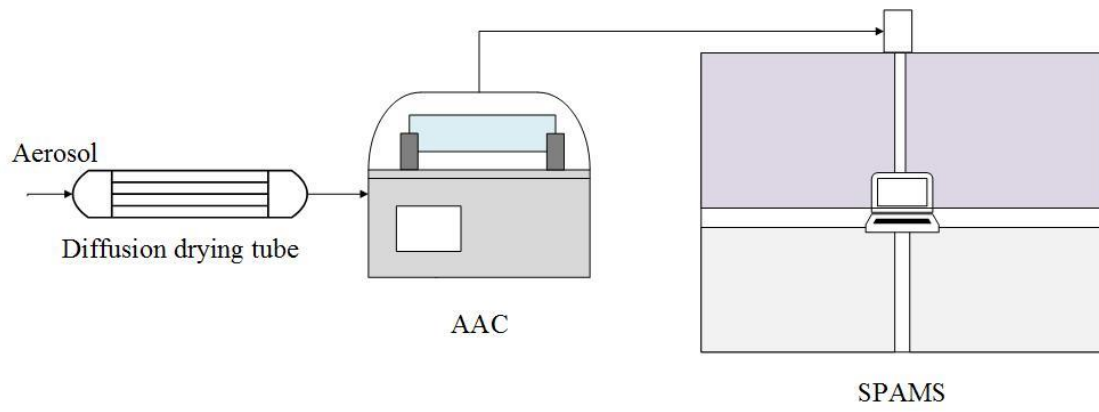
490 Zhang, Y. X., Zhang, Q., Cheng, Y. F., Su, H., Kecorius, S., Wang, Z. B., Wu, Z. J., Hu,
491 M., Zhu, T., Wiedensohler, A., and He, K. B.: Measuring the morphology and
492 density of internally mixed black carbon with SP2 and VTDMA: new insight into
493 the absorption enhancement of black carbon in the atmosphere, *Atmos. Meas.*
494 *Tech.*, 9, 1833-1843, <https://doi.org/10.5194/amt-9-1833-2016>, 2016b.

495 Zhao, G., Zhao, W., and Zhao, C.: Method to measure the size-resolved real part of
496 aerosol refractive index using differential mobility analyzer in tandem with single-
497 particle soot photometer, *Atmos. Meas. Tech.*, 12, 3541-3550,
498 <https://doi.org/10.5194/amt-12-3541-2019>, 2019.

500 **Table 1.** The measured mean D_{ve} and its standard deviation for the eight particle types at D_a values
 501 of 250.0 nm, 350.0 nm, 450.0 nm, and 550.0 nm from nine round measurement.
 502

D_a (nm)	K-rich	EC-S	K-Na	Amine
250.0	193.1 ± 8.2	192.2 ± 8.1	193.8 ± 21.9	190.6 ± 4.6
350.0	284.0 ± 28.4	280.8 ± 9.3	271.9 ± 18.0	284.8 ± 18.2
450.0	364.7 ± 21.1	357.8 ± 6.9	342.5 ± 7.3	367.9 ± 9.7
550.0	416.6 ± 28.3	439.5 ± 5.4	397.3 ± 29.7	442.5 ± 7.4
D_a (nm)	EC-N-S	OC-N-S	OC-EC-N-S	Metal-rich
250.0	188.5 ± 5.9	200.8 ± 17.9	195.4 ± 8.9	189.0 ± 6.7
350.0	281.3 ± 9.3	295.7 ± 29.8	294.0 ± 32.3	277.0 ± 9.1
450.0	358.0 ± 5.8	398.3 ± 44.3	428.9 ± 24.0	342.9 ± 10.0
550.0	453.2 ± 16.4	547.4 ± 14.7	570.9	407.4 ± 14.5

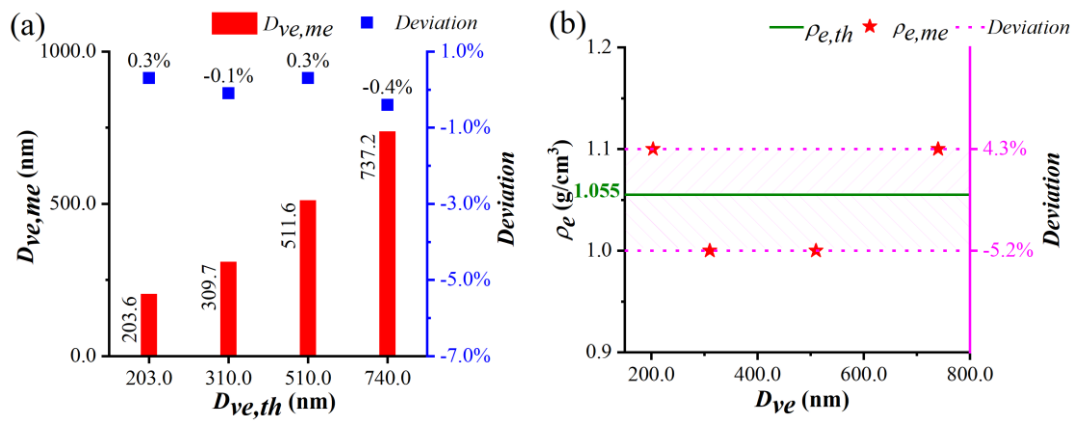
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504

505 **Figure 1.** Schematic diagram of the AAC-SPAMS system (0.3 lpm). The diffusion drying tube is

506 filled with orange silica gel, which reduces the RH to 5-15%.



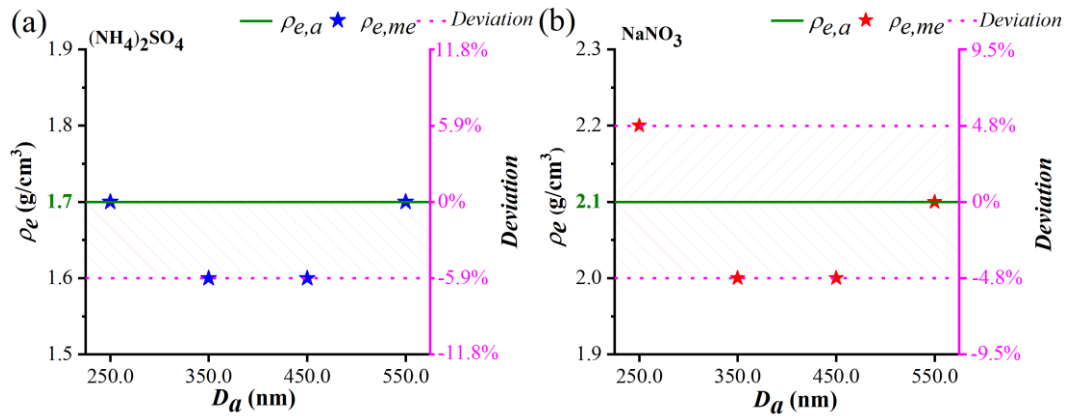
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509 **Figure 2.** (a) Comparison between the measured D_{ve} ($D_{ve,me}$) and the theoretical D_{ve} ($D_{ve,th}$) of the

510 PSL particles. (b) Comparison between the measured ρ_e ($\rho_{e,me}$) and the theoretical ρ_e ($\rho_{e,th}$) of the

511 PSL particles.



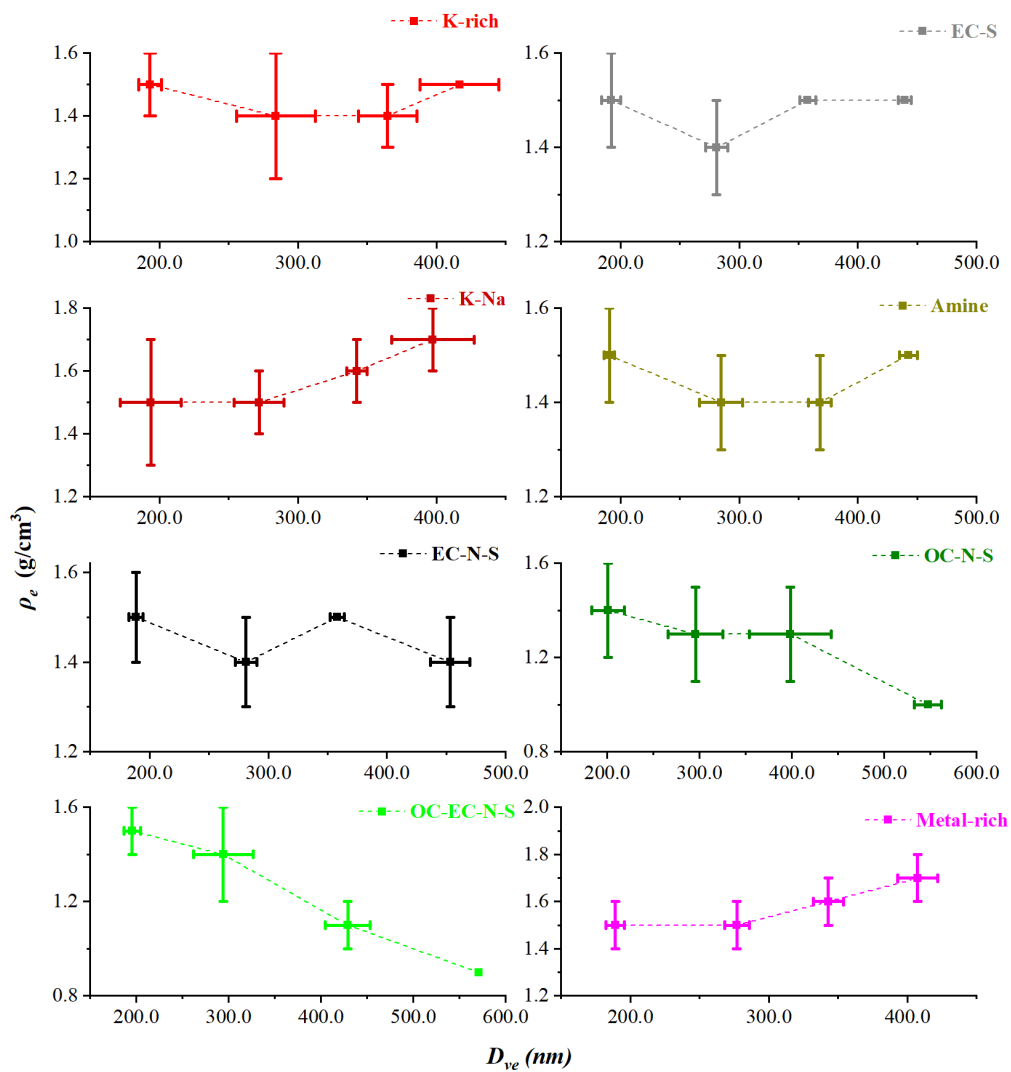
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514 **Figure 3.** (a) Comparison between the measured ρ_e ($\rho_{e,me}$) and average ρ_e ($\rho_{e,a}$) values of the

515 $(\text{NH}_4)_2\text{SO}_4$ particles. (b) Comparison between the measured ρ_e ($\rho_{e,me}$) and average ρ_e ($\rho_{e,a}$) values of

516 the NaNO_3 particles.



517

518

519 **Figure 4.** Variation in ρ_e of the eight particle types with D_{ve} . The solid lines represent the rang of

520 the ρ_e and D_{ve} measured from nine rounds, and the data points stand for the average values.