



1 **Measurement report: Comparison of wintertime individual particles**
2 **at ground level and above the mixed layer in urban Beijing**

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13 **Abstract:**

14 Beijing has been suffering from frequent severe air pollution events, with concentrations affected
15 significantly by the mixed layer height. Major efforts have been made to study the physico-chemical
16 properties, composition, and sources of aerosol particles at ground level. However, little is known
17 on morphology, elemental composition, and mixing state of aerosol particles above the mixed layer.
18 In this work, we collected individual aerosol particles simultaneously at ground level (2 m above
19 ground) and above the mixed layer in urban Beijing (within the Atmospheric Pollution and Human
20 Health in a Chinese Megacity (APHH-Beijing) 2016 winter campaign). The particles were analyzed
21 off-line using transmission electron microscopy coupled with energy dispersive X-ray spectroscopy.
22 Our results showed that the relative number contribution of mineral particles to all measured
23 particles was much higher during non-haze periods (42.5%) than haze periods (18.1%); on the
24 contrary, internally mixed particles contributed more during haze periods (21.9%) than non-haze
25 periods (7.2%) at ground level. In addition, more mineral particles were found at ground level than
26 above the mixed layer height. Around 20% of individual particles showed core-shell structures
27 during haze periods, whereas only a few core-shell particles were observed during non-haze periods
28 (2%). We found that the particle above the mixed layer tend to be more aged with a larger proportion
29 of organic particles originated from coal combustion. Our results indicate that a significant fraction
30 of the airborne particles above the mixed layer originated from surrounding areas influenced by coal
31 combustion activities. This source contributes to the surface particle concentrations in Beijing when
32 polluted air is mixed down to the ground level.

33 **Keywords:** haze; individual particle; organic particle; core-shell structure; mixed layer.



34 **1. Introduction**

35 Atmospheric aerosols, emitted from anthropogenic or natural sources, consist of various
36 chemical constituents (e.g., organic matter, black carbon, nitrate, sulfate, ammonium, metals,
37 mineral dust) (Merikallio et al., 2011; Guo et al., 2014; Wang et al., 2016; Peng et al., 2016; Shao et
38 al., 2017; Tao et al., 2017). Anthropogenic aerosols have received increased attention in the last
39 decades due to their effects on climate and the environment. In fact, anthropogenic aerosols affect
40 climate through cloud condensation nuclei activity (Kerminen et al., 2012), hygroscopic growth
41 (Brock et al., 2016), and light scattering and absorption (Jacobson, 2001; Bond and Bergstrom, 2006;
42 Merikallio et al., 2011; China et al., 2013; Peng et al., 2016; Bhandari et al., 2019b). They can also
43 adversely impact human health; for example, by carrying toxic and carcinogenic compounds (Chen
44 et al., 2013; Shao et al., 2016, 2017). High concentrations of aerosol particles in urban air can cause
45 cardiovascular, respiratory, and even nervous system diseases (Xia et al., 2018; De Marco et al.,
46 2019; Shou et al., 2019). It is suggested that outdoor air pollution causes 3.3 million people
47 premature deaths globally each year (Lelieveld et al., 2015). Atmospheric aerosol particles also
48 affect regional and global geochemical cycles when they are transported over long distances (Heald
49 et al., 2006; Weijun Li et al., 2017; Rodriguez-Navarro et al., 2018).

50 Recently, China has suffered from severe air pollution conditions, like other countries
51 undergoing rapid social and economic development (Huang et al., 2014). In China, this has been
52 associated with frequent occurrence of haze episodes, high PM_{2.5} mass levels, and expanded haze
53 areas (Guo et al., 2014; Huang et al., 2014; Sun et al., 2014). For example, the maximum hourly
54 average PM_{2.5} mass concentrations reached more than 1000 µg m⁻³ in Beijing winter time (Li et al.,
55 2017a; Zhang et al., 2017), 40 times above the safe level of 25 µg m⁻³ recommended by the World
56 Health Organization (WHO).

57 As the megacity capital, Beijing has received much attention being one of the most polluted
58 cities in China. Atmospheric researchers have focused on aerosol particles to understand haze
59 formation in China (Sun et al., 2013; Huang et al., 2014; Zhou et al., 2018b). Measurements and
60 model analyses highlight the key roles of secondary aerosol formation by trace gases (e.g., volatile
61 organic compounds, and SO₂, NO_x) and stagnant meteorological conditions in the regional haze
62 formation (Wang et al., 2013; Guo et al., 2014; Huang et al., 2014).



63 As characterization of aerosol particles has focused on surface level observations, the
64 knowledge of aerosol properties at higher altitudes in urban areas remains poor (Zhou et al., 2018a).
65 Vertical differences between precursors, oxidants and temperature gradients may influence gas-
66 particle partitioning and heterogeneous reactions of N_2O_5 (Zhou et al., 2018a). Previous Beijing
67 measurements at the Institute of Atmospheric Physics (IAP) meteorological tower showed complex
68 vertical distributions of particulate matter and gaseous pollutants (Meng et al., 2008; Sun et al., 2015;
69 Wang et al., 2018; Zhou et al., 2018b), but most of these studies focused on non-refractory
70 submicron species. Research showed that the mixed layer height (MLH) could explain some of the
71 vertical difference of aerosol particle chemical composition (Sun et al., 2015; Wang et al., 2018).
72 For example, vertical distributions of aerosol particles tend to be more uniform during periods with
73 higher MLH (Wang et al., 2018). As heavily increased air pollution can reduce boundary layer
74 heights by diminishing incoming solar energy and therefore, by weakening vertical turbulence,
75 aerosol near-surface concentrations become elevated (Petaja et al., 2016). Moreover, the upper layer
76 particles can influence those below in the MLH by downward entrainment or mixing plumes,
77 making the lower layer particles more complex (Wehner et al., 2010; Platis et al., 2015; Qi et al.,
78 2019). Previous studies showed that the particles above the MLH sometimes considerably influence
79 cloud formation (Camerero et al., 2018) and showed strong aerosol-radiation effect (Bond and
80 Bergstrom, 2006). The differences in aerosol types at ground level and at higher altitudes can lead
81 to large differences in aerosol direct forcing estimates (Ramanathan et al., 2002; Li et al., 2010).
82 The vertical difference of aerosol particles also increases the uncertainties in the assessment of the
83 climate system (Li et al., 2017b). Therefore, a detailed knowledge of the vertical distribution and
84 chemical composition of aerosols is important for understanding the impact on climate and the
85 aerosol evolution process (Zhang et al., 2009; Wang et al., 2018).

86 Vertical comparisons of individual aerosol particles and their morphology, mixing states, and
87 elemental compositions are very limited. Transmission Electron Microscopy (TEM) can provide
88 detailed individual particle characterization and help to explain heterogeneous reactions and aging
89 process (Li et al., 2016a). In this study, we compare particles simultaneously collected at ground
90 level and above the MLH based on the meteorological tower at IAP in Beijing, as part of the UK-
91 CHINA atmospheric pollution and Human Health (APHH) 2016 winter campaign.



92 **2. Experimental**

93 **2.1. Aerosol sampling**

94 Individual aerosol samples were collected at the tower division of IAP, Chinese Academy of
95 Science (39°58'N, 116°22'E), in Beijing from 1 to 9 December of 2016. The site, located between
96 the north 3rd and 4th ring roads in Beijing, is influenced by surrounding and regional traffic, and
97 commercial, as well as, residential activities (Sun et al., 2016).

98 Two DKL-2 single stage cascade impactors, with a 0.5-mm-diameter jet nozzle and a flow rate
99 of 1 L min⁻¹ were used. The sampler collection efficiency is ~ 100% at an aerodynamic diameter of
100 0.5 μm if the particle density is 2 g·cm⁻³ (Li et al., 2016b). Copper (Cu) TEM grids, coated with
101 carbon film (300-mesh; Tianld Co.; Beijing, China), were used to collect the aerosol samples. The
102 sampling duration ranged from 30 second to less than 5 minutes depending on the air pollution loads.
103 Simultaneous observation at ground level (Z1; 2 m above ground) and an elevated altitude (Z2; 280
104 m above ground) allowed us for the vertical profile of the characteristics of the particles to collected.
105 The collected samples were stored in a dry plastic tube and placed in an air dryer to minimize particle
106 changes before analysis.

107 Automatic Lidar and Ceilometer (ACL) observations of attenuated backscatter were conducted
108 at the site using a Vaisala CL31 sensor. Measurements were corrected to account for instrument-
109 related background and near range artefacts (Kotthaus et al., 2016). The MLH was derived from
110 profile measurements using the automatic CABAM algorithm (Kotthaus and Grimmond, 2018).
111 Since the TEM samples were collected for less than 5 minutes, the MLH at 15 min resolution was
112 used to determine whether the Z2 observations were located within the MLH or above the MLH
113 (Shi et al., 2019).

114 Samples were obtained during the periods shown (solid dots and dashed lines) in Fig. 1a;
115 detailed sample information is provided in Table 1. Other measurements including PM_{2.5}, SO₂, NO₂,
116 and O₃ mass concentrations at ground level were obtained from the Olympic Park monitor site,
117 which is the closest national air quality monitor station to IAP (~1.5 km) (Shi et al., 2019). City
118 average temperature (T) and relative humidity (RH) at ground level were obtained from the Ministry
119 of Ecology and Environment of China (<https://www.aqistudy.cn/>).

120 **2.2. Individual particle analysis**



121 Individual particles were analyzed using a JEOL JEM-2800 TEM at an accelerating voltage
122 of 200 kV. The morphology and mixing state of individual particles were determined from the TEM
123 images. Semi-quantitative elemental composition was determined using Energy Dispersive X-ray
124 Spectroscopy (EDS), by which elements heavier than Boron ($Z \geq 6$) can be detected. Cu is not
125 included in this paper because the TEM grids were made of copper. The aerosol particles were not
126 evenly distributed on the TEM grids; the coarser particles occurred near the center and the finer
127 particles occurred on the periphery. To ensure a representative data analysis, three or four areas from
128 the center to the periphery were selected and analyzed. The EDS collection duration of each
129 individual particle was about 15 s to reduce damage of particles from the electron beam. The
130 projected areas of individual particles were determined using the Image-J software (Schneider et al.,
131 2012), which is commonly used for counting and measuring the projected area of atmospheric
132 particles acquired by electron microscopes (Unga et al., 2018). First, the gray-scale images of the
133 particles were converted into binary images, in which black pixels represent the particles and white
134 pixels represent the background. The area equivalent diameters (D_{Aeq}) of the particles were
135 calculated by the following formula: $D_{Aeq} = 2 \cdot (A/\pi)^{1/2}$, where A is the projected area of the particles
136 shown in the TEM image.

137 3. Results and discussions

138 3.1. Air pollutants mass concentrations

139 The temporal variations of different air pollutants and meteorological conditions at ground
140 level are shown in Fig. 1. The hourly averaged $PM_{2.5}$ mass concentration at the Olympic Park
141 monitoring site ranged from 3 to 530 $\mu\text{g m}^{-3}$, with a sample period average of 113.3 $\mu\text{g m}^{-3}$,
142 significantly exceeding the safe level of 75 $\mu\text{g m}^{-3}$ according to the Chinese National Ambient Air
143 Quality Standard (GB3095-2012). The MLH ranged from 54 to 1496 m, with an average of 397 m.
144 The MLH showed obvious daily cycles. The hourly mean RH ranged from 17% to 97%, with a 9
145 day mean of 50.3%. The RH and $PM_{2.5}$ are positively correlated (correlation coefficient=0.75; Fig.
146 S1) according to the 216 groups of hourly data, suggesting that the higher RH favors the formation
147 of haze (Sun et al., 2014; Wang et al., 2016). As expected, RH and temperature were negatively
148 correlated (correlation coefficient=-0.51; Fig. S2). The SO_2 time series has similar trends to that of
149 NO_2 . However, the average concentration of NO_2 (83.2 $\mu\text{g m}^{-3}$) was nearly 5.5 times higher than



150 that of SO₂ (15.2 μg m⁻³). The concentration of O₃ showed a different trend compared with NO₂ and
151 SO₂ (Fig. 1), with a 9 days hourly mean concentration of 20 μg m⁻³.

152 **3.2. Classification and mixing state of individual particles**

153 Aerosol particles were classified using their morphologies and elemental compositions into
154 seven main types, namely: 1) organic particles (OPs), 2) sulfur-rich (S-rich) particles, 3) soot
155 particles, 4) mineral particles, 5) metal particles, 6) internally mixed organic and sulfur-rich particles
156 (OP-S), and 7) other mixed particles. The detailed characteristics of each particle type are shown in
157 Table 2.

158 OPs are mainly composed of C and O, usually with a small amount of Si, S, Cl, and K. OPs
159 are relatively stable under the electron beam irradiation. Based on the morphologies, OPs can be
160 further divided into spherical (Fig. 2a) and irregularly shaped (Fig. 2b). They were mainly from
161 combustion process of biomass and fossil fuel.

162 S-rich particles (Figs. 2c and 2d) are mainly composed of O, S, and N, and sometimes also
163 contain some amount of K. S-rich particles are beam-sensitive and volatilize under strong beam
164 irradiation. S-rich particles generally represent secondary inorganic components (e.g., SO₄²⁺, NO₃⁻
165 and NH₄⁺) (Xu et al., 2019).

166 Soot particles are mainly composed of C, minor amount of O, and sometimes Si. Soot particles
167 consist of a number of C-dominated spherical monomers less than 100 nm in diameter (Figs. 2e and
168 2f) and can be easily identified under high-resolution TEM (Buseck et al., 2014; Bhandari et al.,
169 2017). Soot particles, stable under the electron beam, show chain-like or compact morphology in
170 the atmosphere (Sorensen et al., 2001; Adachi et al., 2007; China et al., 2013, 2015; Bhandari et al.,
171 2019a). Soot particles are mainly generated during incomplete combustion of biomass and fossil
172 fuel.

173 Metal particles (Figs. 2g and 2h) and Mineral particles (Fig. 2i) are stable under the beam
174 irradiation. Mineral particles are mostly irregularly shaped containing crustal elements (e.g., Si, Al,
175 Ca, Fe, Na, K, Mg, Ti, and S) in addition to O. They can be generated from windblown soil dust or
176 road dust. Metal particles are spherical or near spherical and are mainly composed of Fe, Zn, Mn,
177 Ti, and Pb. Metal particles are mainly originating from industries, coal-fired power plants, and oil
178 refineries (Xu et al., 2019) or vehicle brakes (Hou et al., 2018).



179 Internally mixed particles (Figs. 2j-p) are particles with at least two of the above components.
180 They usually show relative larger diameter. We further classify them as internally mixed organic
181 and sulfur-rich particles (OP-S) (Figs. 2j-l), and other mixed particles (Figs. 2m-p).

182 3.3. Ground level haze and non-haze individual particle comparison

183 Haze periods are defined as when the hourly average $PM_{2.5}$ mass concentration is more than
184 $75 \mu g m^{-3}$; the rest are defined as non-haze periods. A total of 1538 individual particles among 8
185 samples at ground level were analyzed based on the TEM results. The relative number percentage
186 $(N(\text{type } i)/N(\text{total}) \times 100)$ of particles in each sample was calculated, and the results are provided in
187 Table 3 and shown in Fig.3. During non-haze periods, the particles were composed of mineral
188 particles (42.5%), OPs (21.1%), S-rich particles (20.0%), soot particles (6.4%), other mixed
189 particles (5.6%), metal particles (2.83%), and OP-S (1.6%) in descending order. During haze periods,
190 the particles were composed of OPs (28.3%), S-rich particles (23.5%), mineral particles (18.1%),
191 OP-S (13.1%), other mixed particles (8.8%), soot particles (6.6%), and metal particles (1.7%) in
192 descending order.

193 The mineral particles are mainly from re-suspended road dust, soil dust, and construction dust
194 during non-desert transport dust episodes (Sun et al., 2006; Gao et al., 2016; Wang et al., 2017). The
195 relative number percentage of mineral particles was much higher during non-haze periods (42.5%)
196 than during haze periods (18.1%), as shown in Fig.3. However, the mixed particles including OP-S
197 and other mixed particles were much more abundant during haze periods (21.9%) than during non-
198 haze periods (7.2%), suggesting that there was more secondary aerosol formation during haze
199 periods. High secondary aerosol formation in winter in Beijing during the pollution periods was also
200 found in previous studies (Huang et al., 2014; Sun et al., 2016; Li et al., 2017a). Secondary aerosol
201 formation is expected since the RH during the haze periods were relatively higher than during non-
202 haze periods, as shown in Table 1 and Fig.1, which facilitated chemical reactions of gaseous
203 pollutants (Liu et al., 2016; Wang et al., 2016). Also, the average OPs and S-rich were higher during
204 haze periods than during non-haze periods.

205 3.4. Ground level and above the MLH individual particle comparison

206 A total of 1519 individual particles from 8 samples above the MLH were analyzed. The results
207 are provided in Table 3 and shown in Fig. 3. During non-haze periods, the contribution of mineral



208 particles above the MLH (23.2%) was less than that at ground level (42.5%), but the S-rich and OPs
209 accounted for 30.7% and 27.3% above the MLH, respectively, fractions higher than those of 20.0%
210 and 21.1% at the ground level. During haze periods, the contribution of mineral particles above the
211 MLH (9.5%) was also lower than at ground level (18.1%). S-rich particles were also less abundant
212 above the MLH (16.4%) compared to the ground level (23.5%), which is different from the non-
213 haze periods. This may be because more S-rich particles above the MLH were mixed with other
214 particles, forming mixed particles. The mixed particles above the MLH were much higher than at
215 ground level, especially the OP-S particles (20.7% vs 13.1%). OPs above the MLH (34.8%) were
216 more abundant than at ground level (28.5%). Particles above the MLH were either transported from
217 the surrounding areas or from ground sources. In both cases, they were subject to atmospheric
218 process, leading to their aging.

219 3.5. Aging of particles

220 During the aging process of aerosol particles, secondary species can coat pre-existing particles
221 (Li and Shao, 2009; Laskin et al., 2016; Li et al., 2016b; Niu et al., 2016; Tang et al., 2016; Chen et
222 al., 2017; Hou et al., 2018; Unga et al., 2018; Xu et al., 2019). Using high-resolution TEM images,
223 it is possible to identify the core-shell structure of particles (Li et al., 2016a). For example, Figs. 4a
224 and 4b show S-rich particles coated by secondary species. Figs. 4c and 4d show Ops that were
225 coated with secondary species. Figs. 4e-h show core-shell structured particles with some mixed
226 particle cores. In this study, we found that the core-shell structured particles accounted for 20%
227 during haze periods but only 2% during non-haze periods. Also, the average D_{Aeq} of particles was
228 larger during haze periods than during non-haze periods as shown in Fig. S3. These results
229 confirmed that particles during haze periods underwent more extensive aging than during non-haze
230 periods.

231 The coating of atmospheric particles is often caused by aging mechanisms such as coagulation,
232 condensation, and heterogeneous chemical reactions (Kahnert, 2015; Müller et al., 2017). Fig. 5
233 shows low magnification images of particles at ground level and above the MLH. The core/shell
234 ratio (R), which is the ratio of the D_{Aeq} of the core to the D_{Aeq} of the whole particle including the
235 coating, has been used to evaluate the aging process of aerosol particles in different studies (Niu et
236 al., 2012, 2016; Hou et al., 2018). The value of R ranged from 0 to less than 1. A smaller R value



237 means the particles were more coated, thus were subjected to a more extensive degree of aging (Hou
238 et al., 2018). Because a high number percentage of core-shell structured particles were only found
239 during haze periods, we measured R of core-shell structured particles only during the haze periods
240 (including the samples 2, 4, 5, 6 and 7). Fig. 6a shows the R value of each samples during the haze
241 periods. The average R value above the MLH (0.54) was smaller than ground level (0.59). We can
242 see from Fig. 6a that all the samples showed a smaller average R value above the MLH compared
243 with those from the ground level. Additionally, the relative number percentage of core-shell
244 structured particles was higher above the MLH than at ground level, except for sample 4. These
245 findings indicate that the particles above the MLH were more aged than those at ground level.

246 **4. Summary and Atmospheric implications**

247 Our results show that mineral particles represented a higher number percentage during non-
248 haze periods (42.5%) than during haze periods (18.1%), and mixed particles were more abundant in
249 haze periods (21.9%) than in non-haze periods (7.2%) at the ground level. In addition, more mineral
250 particles were found at ground level than above the MLH. Our results also show higher relative
251 number percentage of OPs both during non-haze (21.1%) and haze periods (28.3%) in winter
252 Beijing, compared with a tunnel environment (~5%), where the vehicle emissions were the main
253 pollution sources (Hou et al., 2018). Also, recent studies did not find abundant OPs in North China
254 during Spring and Summer (Yuan et al., 2015; Li et al., 2016b; Xu et al., 2019;). Instead, a larger
255 number percentage of OPs have been found in winter using electron microscopy in previous studies,
256 including an outflow of a haze plume in East Asia (Zhu et al., 2013), a coal-burning region in China's
257 Loess Plateau (Li et al., 2012), three sampling sites in North China Plain (Chen et al., 2017) and
258 urban and rural sites in Northeast China (Xu et al., 2017; Zhang et al., 2017). The results above
259 suggest that OPs account for a large number percentage of the particles in north China in winter.

260 Most of the OPs in our study were spherical or nearly spherical in shape according to the
261 projected images, suggesting that they were formed through cooling process after the biomass or
262 fossil fuel combustion pyrolysis products of volatile organic compounds were emitted into the
263 atmosphere (Wang et al., 2015; Chen et al., 2017; Zhang et al., 2017). These spherical or near
264 spherical OPs were considered to be brown carbon (Zhang et al., 2020). Brown carbon plays a
265 significant role in atmospheric shortwave absorption and can cause warming of the atmosphere



266 (Adachi and Buseck, 2011; Hoffer et al., 2016). Some researchers have found that the primary OPs
267 from coal combustion has more Si than those from biomass burning (Li et al., 2012; Chen et al.,
268 2017). The weight ratio of C-O-Si at ground level and above the MLH is shown in Fig. 7. More coal
269 burning related OPs were found above the MLH. Since the relative number percentage of primary
270 OPs affected by coal burning are higher above the MLH than at the ground level, the OPs above the
271 MLH are not all from the ground level and might have originated from surrounding areas influenced
272 by coal combustion. The particles above the MLH can contribute to Beijing air pollution by mixing
273 down to the ground.

274 In this study, more core-shell structured particles were found above the MLH than at the
275 ground; finding which can have important atmospheric implications. Fig. S3 shows the total particle
276 number-size distribution; the larger size particles clearly increased when considering the coatings
277 compared to only considering the core size during haze periods. The changes in optical properties
278 due to coating was calculated in various studies by using different methods (Cappa et al., 2012;
279 Scarnato et al., 2013; Liu et al., 2015; Saliba et al., 2016; Unga et al., 2018). When host particles
280 are coated, their optical properties might be amplified. For example, when soot particles (optically
281 often referred to as black carbon) are coated by secondary species, the light absorption is typically
282 enhanced because of the so called “lensing effect” (Khalizov et al., 2009; Liu et al., 2009; Peng et
283 al., 2016). Previous measurements showed that soot can heat the upper boundary layer more than
284 the lower layer during haze periods and decrease surface heat flux substantially, thus depressing the
285 development of MLH (Ding et al., 2016). Also, organic coating can influence the hygroscopic
286 properties and the viscosity of mixed particles (Sharma et al., 2018; Unga et al., 2018), and thus can
287 influence cloud formation activity (Kerminen et al., 2012).

288 The different relative number percentage of particle types and different aging degree of the
289 particles have important implications for understanding the climate effects of aerosol particles and
290 for emission control policy making.

291 **Data availability:** Data used in this study are available from the corresponding author upon request
292 (ShaoL@cumtb.edu.cn)

293 **Author Contributions:** WW, LS, CM, JX and ZS conceived the manuscript. WW, WL, XF and
294 MZ conducted the sample collection and analysis. SK and SG conducted the MLH measurement.
295 CM and BJ conducted manuscript modification.



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- 596



597 Table 1 Sample information and meteorological conditions

Sample ID ^①	Date (2016)	Time ^②	PM _{2.5} (µg m ⁻³)	SO ₂ (µg m ⁻³)	NO ₂ (µg m ⁻³)	O ₃ (µg m ⁻³)	RH (%)	T (°C)	MLH (m) ^③
Z1-1	12/1	9:10	12 ^④	2	48	37	24	6	--
Z2-1	12/1	8:40	--	--	--	--	--	--	194
Z1-2	12/2	1:00	110	25	109	3	55	2	--
Z2-2	12/2	1:00	--	--	--	--	--	--	141
Z1-3	12/2	9:10	24	20	134	2	50	3	--
Z2-3	12/2	8:40	--	--	--	--	--	--	134
Z1-4	12/3	1:53	142	36	102	6	79	-1	--
Z2-4	12/3	3:00	--	--	--	--	--	--	232
Z1-5	12/4	1:04	530	14	180	4	93	1	--
Z2-5	12/4	3:00	--	--	--	--	--	--	136
Z1-6	12/5	2:00	86	8	21	53	75	2	--
Z2-6	12/5	2:00	--	--	--	--	--	--	114
Z1-7	12/8	9:10	187	2	16	72	86	2	--
Z2-7	12/8	8:40	--	--	--	--	--	--	191
Z1-8	12/9	9:20	12	8	67	12	33	2	--
Z2-8	12/9	8:30	--	--	--	--	--	--	250

598 ^①Samples were collected at two altitudes: Z1 was 2 m above ground and Z2 was 280 m above
599 ground. ^② Sampling duration ranged from 30 s to less than 5 minutes depending on the PM
600 pollution. ^③ MLH represents the mixed layer height and the data are 15minutes average; MLH
601 was less than 280 m and the samples collected at Z2 represent samples above the mixed layer. ^④
602 If PM_{2.5} mass concentration was less than 75 µg m⁻³, samples were classified as non-haze samples
603 and if PM_{2.5} mass concentration was more than 75 µg m⁻³, samples were classified as haze samples.
604



605 Table 2 Classification and characteristics of individual particle types.

Particle type	Elemental composition	Morphology	Possible sources
Soot particles	C and minor amounts of O, Si.	Chain-like or compact C-dominated aggregates.	Incomplete combustion of biomass and fossil fuel.
Organic particles	C and O with minor amounts of Si, K, S, Cl.	Spherical, near spherical or irregular shapes.	Combustion process or secondary aerosol formation.
Mineral particles	O, Si, Al, Ca, Fe, Na, K, Mg, Ti, and S.	Irregular shapes.	Re-suspended from soil dust, road dust, and construction dust.
Metal particles	Fe, Zn, Mn, Ti, and Pb.	Spherical or irregular shapes.	Industries, coal-fired power plants and oil refineries.
S-rich particles	S and O with minor amounts of N, K.	Spherical, near spherical or irregular shapes.	Secondary aerosol formation.
Organic mixed with Sulfur-rich particles	C, O, and S with minor amounts of N, K or Cl.	Irregular shapes.	Secondary aerosol reaction.
Other mixed particles	Complex elemental composition.	Irregular shapes with different particle types.	Secondary aerosol reaction.

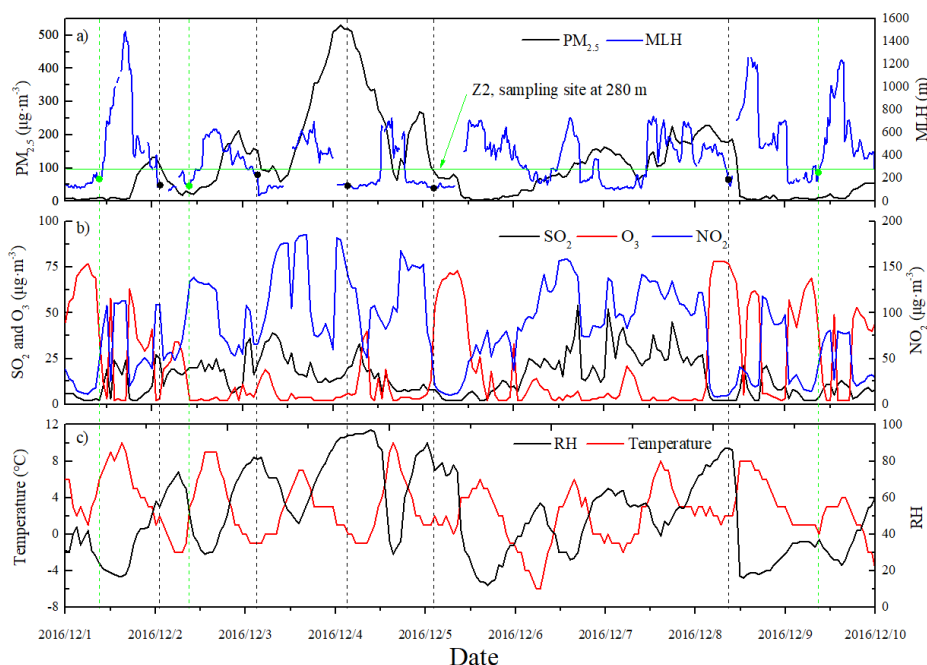
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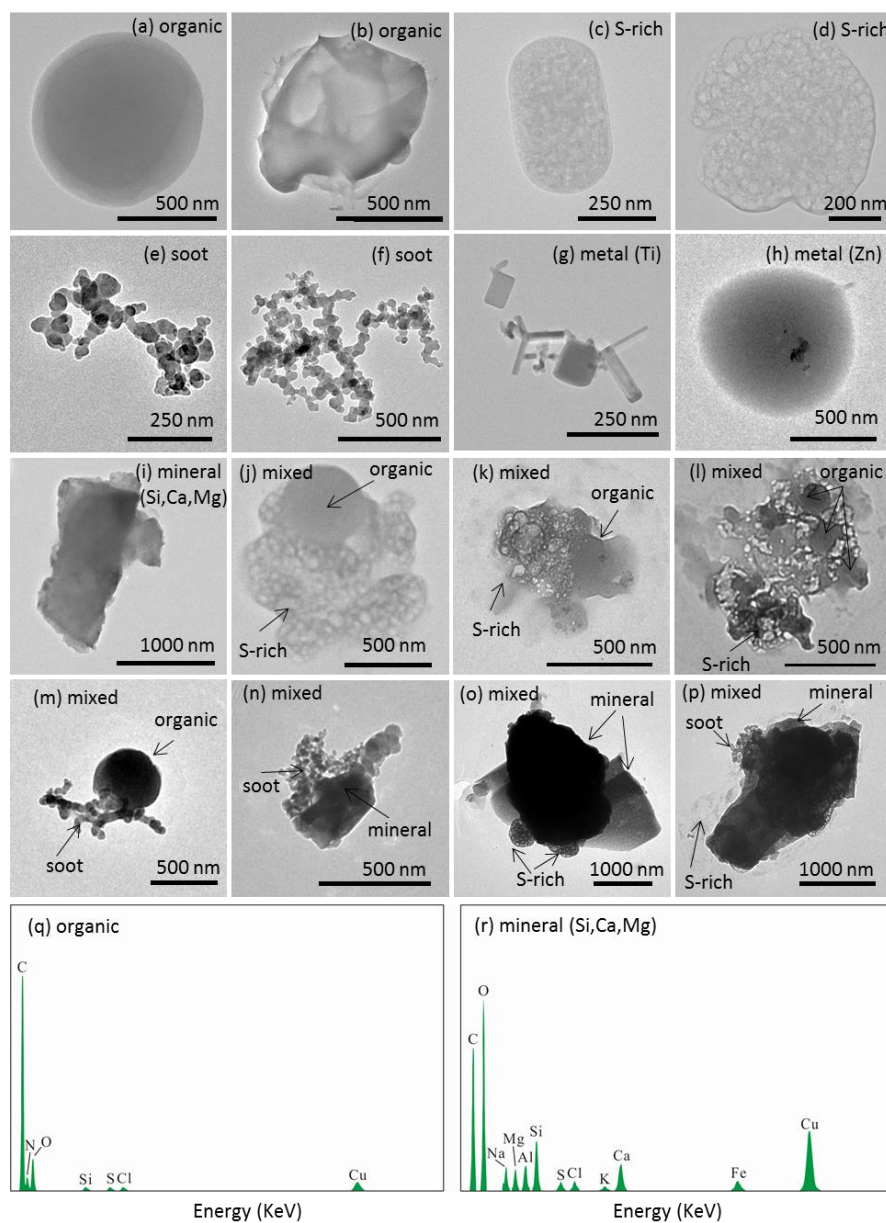
607 Table 3 Relative number percentage of individual particles.

Air qualities	Sample ID	Number	Metals	Minerals	OPs	S-rich	Soot	OP-S	Other
Non-haze periods	Z1-1	114	2.6	30.7	19.3	36.0	5.3	1.8	4.4
	Z2-1	113	1.8	12.4	16.8	56.6	10.6	0.9	0.9
	Z1-3	135	4.4	34.1	31.9	12.6	11.1	0.7	5.2
	Z2-3	118	2.5	23.7	45.8	17.0	4.2	2.5	4.2
	Z1-8	140	1.4	62.9	12.1	11.4	2.9	2.1	7.1
	Z2-8	119	3.4	33.6	19.3	18.5	17.7	0.0	7.6
	Ave (Z1)	389	2.8	42.5	21.1	20.0	6.4	1.6	5.6
	Ave (Z2)	350	2.6	23.2	27.3	30.7	10.8	1.1	4.2
Haze periods	Z1-2	123	2.4	21.1	42.3	17.1	7.3	2.4	7.3
	Z2-2	164	4.9	14.6	37.2	25.0	4.3	9.8	4.3
	Z1-4	160	0.6	28.8	30.6	8.8	13.8	9.4	8.1
	Z2-4	266	0.0	3.8	53.0	3.4	7.1	19.6	13.2
	Z1-5	461	0.9	6.5	18.9	22.1	7.6	31.5	12.6
	Z2-5	266	0.4	0.4	32.3	7.1	2.3	44.0	13.5
	Z1-6	237	2.5	11.0	21.5	48.5	2.1	6.8	7.6
	Z2-6	281	1.8	11.0	18.9	19.6	12.8	15.3	20.6
	Z1-7	168	1.8	23.2	28.0	20.8	2.4	15.5	8.3
	Z2-7	192	1.6	17.7	32.3	27.1	1.6	15.1	4.7
	Ave (Z1)	1149	1.7	18.1	28.3	23.5	6.6	13.1	8.8
	Ave (Z2)	1169	1.7	9.5	34.7	16.4	5.6	20.7	11.3

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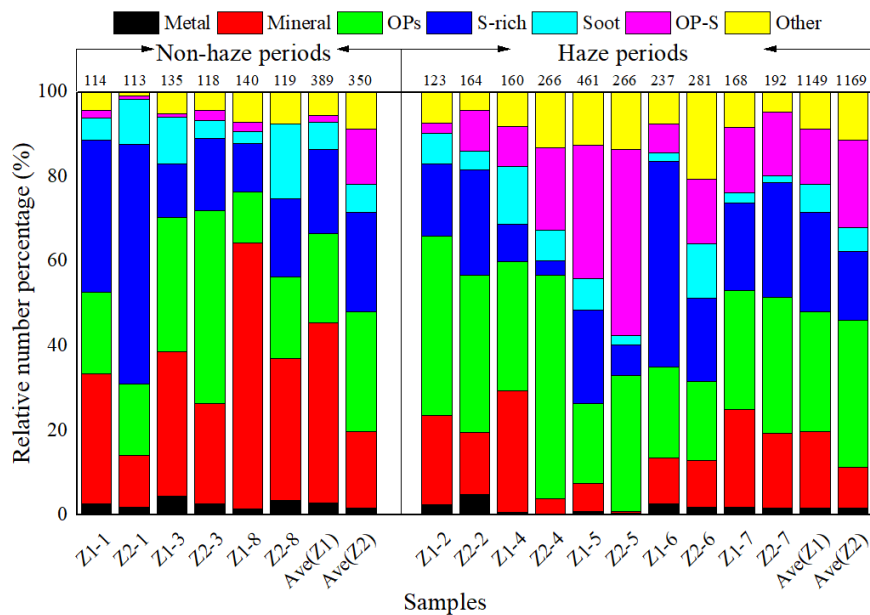
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610 Fig. 1: The dashed lines represent the individual particle sampling times with green lines
611 representing non-haze samples and black lines haze samples. (a) Temporal variations of mixed layer
612 height (MLH) and $PM_{2.5}$ mass concentrations. The solid dots represent the MLH during the
613 sampling times. (b) Temporal variations of SO_2 , NO_2 , O_3 at ground level at the Olympic Park
614 monitor site, which is the closest national air quality monitor station to the sampling site (~1.5 km).
615 (c) Temporal variations of temperature (T) and relative humidity (RH) at ground level. Date were
616 obtained from the Ministry of Ecology and Environment of China (<https://www.aqistudy.cn>);
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619 Fig. 2: Examples of morphologies and mixing characteristics of individual aerosol particles in
620 winter in Beijing at ground level and above the mixed layer. (a) Spherical organic particle, (b)
621 irregular shaped organic particle, (c-d) S-rich particles, (e-f) soot particles, (g-h) metal particles, (i)
622 mineral particles, (j-l) OP-S mixed particles, and (m-p) other mixed particle types. (q) and (r) are
623 EDS of (b) and (i). The difference between the particles in (b) and (i) is that organic particles (b)
624 mainly composed C and O while minerals (i) mainly composed O, Si, Ca and Mg.
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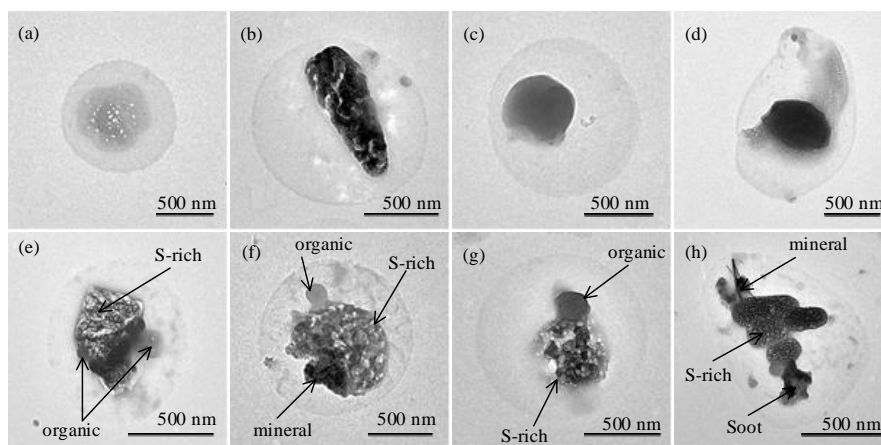


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628 Fig. 3: Relative number percentage of different particle types at ground level (Z1) and above the
 629 mixed layer height (Z2). The number above each bar represents the total particle number analyzed
 630 in each sample.



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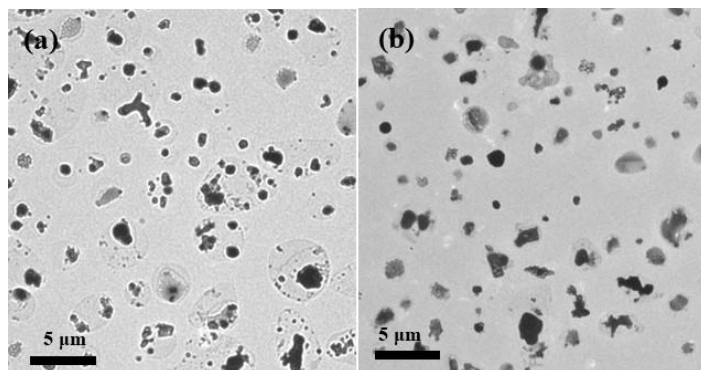


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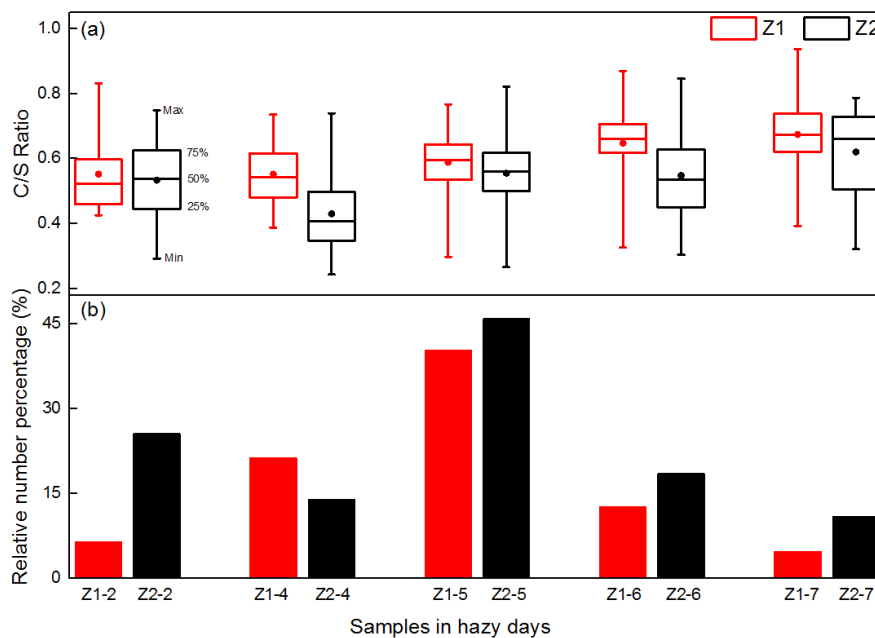
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Fig. 4: Images of core-shell structured particles. (a-b) S-rich cores, (c-d) organic cores, and (e-h) mixed cores.

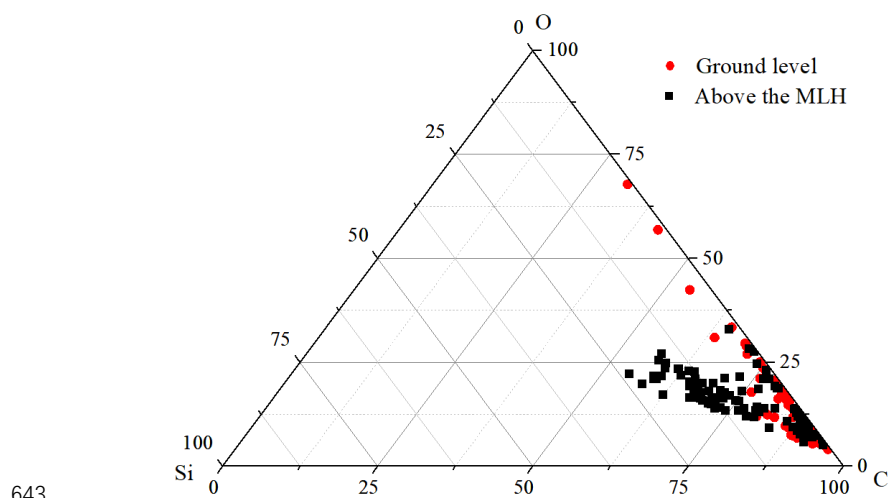


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636 Fig. 5 Low magnification images of individual particles during haze periods above the MLH (a)
637 and at ground level (b). More coated particles are shown above the MLH.



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Fig. 6: (a) C/S ratio (D_{Aeq} ratio of the core to the whole particle including the shell) of particles during haze periods at ground level (Z1) and above the mixed layer height (Z2); solid dots represent the average value, and (b) the corresponding relative number percentage of core-shell structured particles.



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Fig. 7: Triangular diagram showing the weight ratio of C-O-Si of organic particles at ground level and above the mixed layer height (MLH).