



1 **Changes in black carbon emissions over Europe due to COVID-19**

2 **lockdowns**

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43 **Abstract**

44 Following the emergence of the severe acute respiratory syndrome coronavirus 2 (SARS-
45 CoV-2) responsible for COVID-19 in December 2019 in Wuhan (China) and its spread to the
46 rest of the world, the World Health Organization declared a global pandemic in March 2020.
47 Without effective treatment in the initial pandemic phase, social distancing and mandatory
48 quarantines were introduced as the only available preventative measure. Despite the
49 socioeconomic impacts, air quality improved due to lower pollutant emissions. Here we
50 investigate the effects of the COVID-19 lockdowns on ambient black carbon (BC), which
51 affects climate and damages health, using in-situ observations from 17 European stations in a
52 Bayesian inversion framework. BC emissions declined by 11% in Europe (20% in Italy, 32%
53 in Germany, 20% in Spain) during lockdown compared to the same period in the previous five
54 years. BC temporal variation in the countries enduring the most drastic restrictions showed the
55 most distinct lockdown impacts. Increased particle light absorption in at the beginning of the
56 lockdown, confirmed by assimilated satellite and remote sensing data, suggests residential
57 combustion was the dominant BC source. Accordingly, in Central and Eastern Europe, which
58 experienced lower than average temperatures, BC was elevated compared to the previous five
59 years. Except for the comparison of BC emissions in the lockdown with the previous five years,
60 an immediate decrease was also seen, as compared with the period before the lockdown, which
61 averaged about 10% over Europe. Such a decrease was not seen in the previous years, which
62 also confirms an impact on BC emissions from COVID-19.

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65 **1 Introduction**

66 The identification of the severe acute respiratory syndrome coronavirus 2 (SARS-CoV-2
67 or COVID-19) in December 2019 (WHO, 2020) in Wuhan (China) and its subsequent
68 transmission to South Korea, Japan, and Europe (initially mainly Italy, France and Spain) and
69 the rest of the world led the World Health Organization to declare a global pandemic by March
70 2020 (Sohrabi et al., 2020). Although the symptoms are normally mild or not even detected for
71 most of the population, people with underlying diseases or elderly are very vulnerable showing
72 complications that can lead to death (Huang et al., 2020). Considering the lack of available
73 treatment and vaccination to combat further spread of the virus, the only prevention measures
74 included strict social, travel and working restrictions in a so-called lockdown period that lasted
75 for several weeks (mid-March to end of April 2020 for most of Europe). The most drastic
76 measures were taken in China, where the outbreak started, in Italy that faced large human losses
77 and later in the United States. Despite all these restriction, still six months after the first
78 lockdown, several countries are reporting severe human losses due to the virus (John Hopkins
79 University of Medicine, 2020).

80 Despite the dramatic health and socioeconomic consequences of COVID-19 lockdowns,
81 their environmental impact might be beneficial. Bans on mass gatherings, mandatory school
82 closures, and home confinement (He et al., 2020; Le Quéré et al., 2020) during lockdowns have
83 all resulted in lower traffic-related pollutant emissions and improved air quality in Asia, Europe
84 and America (Adams, 2020; Bauwens et al., 2020; Berman and Ebisu, 2020; Conticini et al.,
85 2020; Dantas et al., 2020; Dutheil et al., 2020; He et al., 2020; Kerimray et al., 2020; Le et al.,
86 2020; Lian et al., 2020; Otmani et al., 2020; Sicard et al., 2020; Zheng et al., 2020). The
87 restrictions also present an opportunity to evaluate the cascading responses from the interaction
88 of humans, ecosystems, and climate with the global economy (Diffenbaugh et al., 2020).

89 Strongly light absorbing black carbon (BC, or ‘soot’), is produced from incomplete
90 combustion of carbonaceous fuels e.g. fossil fuels, wood burning, biofuels (Bond et al., 2013).
91 By absorbing solar radiation, it warms the air, reduces tropical cloudiness (Ackerman, 2000)
92 and atmospheric visibility (Jinhuan and Liquan, 2000). BC causes pulmonary diseases (Wang
93 et al., 2014a), may act as cloud condensation nuclei affecting cloud formation and precipitation
94 (Wang et al., 2016) and contributes to global warming (Bond et al., 2013; Myhre et al., 2013;
95 Wang et al., 2014a). When deposited on snow, it reduces snow albedo (Clarke and Noone,
96 1985; Hegg et al., 2009) accelerating melting. Since BC is both climate relevant and strongly



97 linked to anthropogenic activity, it is important to determine the effects of the COVID-19
98 lockdowns thereon.

99 Here, we present a rigorous assessment of temporal and spatial changes BC emissions
100 over Europe (including Middle East and parts of North Africa), combining in situ observations
101 from the Aerosol, Clouds and Trace Gases Research Infrastructure (ACTRIS) network and
102 state-of-the-art emission inventories within a Bayesian inversion. We validate our results with
103 independent satellite data and compare them to inventories and baseline and optimized
104 emissions calculated for previous years.

105 **2 Methods**

106 **2.1 Particle light absorption measurements**

107 The measurement sites contributing data to this paper are regional background sites
108 (except for one site in Germany) and all contribute to the research infrastructure Aerosol,
109 Clouds and Trace Gases Research Infrastructure (ACTRIS), and the networks European
110 Monitoring and Evaluation Program (EMEP) and Global Atmosphere Watch (GAW). The
111 measurement data used for the period 2015 - May 2020 consist of hourly-averaged, quality-
112 checked, particle light absorption measurements. The quality assurance and quality control
113 correspond to the Level 2 requirements for ACTRIS, EMEP and GAW data, as described in
114 detail in Laj et al. (2020).

115 All absorption measurements within ACTRIS and EMEP are taken using a variety of
116 filter-based photometers: Multi-Angle Absorption Photometers (MAAP), Particle Soot
117 Absorption Photometers (PSAP) Continuous Light Absorption Photometers (CLAP), and the
118 Aethalometer (AE31). Information on instrument type at the various sites are included in Table
119 1 and procedures for harmonization of measurement protocols to produce comparable data sets
120 are described in Laj et al. (2020) in detail. Zanatta et al. (2016) suggested that a MAC value of
121 $10 \text{ m}^2 \text{ g}^{-1}$ (geometric standard deviation of 1.33) at a wavelength of 637 nm can be considered
122 to be representative of the mixed boundary layer at European ACTRIS background sites, where
123 BC is expected to be internally mixed to a large extent. Assuming an absorption Ångström
124 exponent (AAE) is equal to unity, i.e. assuming no change in MAC for different sources (Zotter
125 et al., 2017), we extrapolated the MACs at 637 nm ($MAC_{@637}$) to the measurement wavelengths
126 of our study ($MAC_{@λ_2}$) using the following equation:



127 $MAC_{@λ2} = MAC_{@λ1} \left(\frac{λ1}{λ2}\right)^{AAE} \xrightarrow{yields} MAC_{@λ2} = 10 \left(\frac{637}{λ2}\right)^1$ (1)
128 following Lack and Langridge (2013). The resulting MAC values for each measurement
129 station are shown in Table 1.

130 2.2 Source – receptor matrix (SRM) calculations

131 Source – receptor matrices (SRMs), also known as “footprint emission sensitivities” or
132 “footprints” for each of the 17 receptor sites (Table 1) were calculated using the Lagrangian
133 particle dispersion model FLEXPART version 10.4 (Pisso et al., 2019). The model releases
134 computational particles that are tracked backward in time based on 3-hourly operational
135 meteorological analyses from the European Centre for Medium-Range Weather Forecasts
136 (ECMWF) with 137 vertical layers and a horizontal resolution of $1^\circ \times 1^\circ$. The tracking of BC
137 particles includes gravitational settling for spherical particles with an aerosol mean diameter of
138 $0.25 \mu\text{m}$ and a logarithmic standard deviation of 0.3 and a particle density of 1500 kg m^{-3} (Long
139 et al., 2013). FLEXPART also simulates dry and wet deposition (Grythe et al., 2017),
140 turbulence (Cassiani et al., 2014), unresolved mesoscale motions (Stohl et al., 2005) and
141 includes a deep convection scheme (Forster et al., 2007). SRMs were calculated for 30 days
142 backward in time, at temporal intervals that matched measurements at each receptor site. This
143 backward tracking is sufficiently long to include almost all BC sources that contribute to surface
144 concentrations at the receptors given a typical atmospheric lifetime of 3–11 days (Bond et al.,
145 2013).

146 2.3 Bayesian inverse modelling

147 The Bayesian inversion framework FLEXINVERT+ described in detail in Thompson
148 and Stohl (2014) was used to optimize emissions of BC before (January to mid-March) and
149 during the COVID-19 lockdown period in Europe (mid-March to end of April). The algorithm
150 finds the optimal emissions, which lead to FLEXPART-modelled concentrations that better
151 match the observations considering the uncertainties for observations, prior emissions and
152 SRMs. The optimized (a posteriori) emissions are those that minimize the cost function:

$$153 \quad J(\mathbf{x}) = \frac{1}{2}(\mathbf{x} - \mathbf{x}_b)^T \mathbf{B}^{-1}(\mathbf{x} - \mathbf{x}_b) + \frac{1}{2}(\mathbf{y} - \mathbf{H}\mathbf{x})^T \mathbf{R}^{-1}(\mathbf{y} - \mathbf{H}\mathbf{x}) \quad (2)$$

154 where \mathbf{H} is the Jacobian matrix of SRMs, \mathbf{y} the vector of observed BC concentrations, \mathbf{x} and
155 \mathbf{x}_b the vectors of optimized and prior emissions, respectively, while \mathbf{B} and \mathbf{R} are the error
156 covariance matrices for the prior emissions and the observations, respectively. The most
157 probable posterior emissions, \mathbf{x} are given by the following equation (Tarantola, 2005):

$$158 \quad \mathbf{x} = \mathbf{x}_b + \mathbf{B}\mathbf{H}^T(\mathbf{H}\mathbf{B}\mathbf{H}^T + \mathbf{R})^{-1}(\mathbf{y} - \mathbf{H}\mathbf{x}_b) \quad (3)$$



159 Here, posterior emissions were calculated weekly between 1 January and 30 April 2020.
160 To show potential differences on the signal from the 2020 restrictions, we calculated the
161 posterior emissions of BC for the same periods in the years 2015–2019. The aggregated
162 inversion grid (25°N–75°N and 10°W–50°E) and the average SRM for inversions are shown in
163 Figure 1, while the measurement stations are listed in Table 1. The variable grid uses high
164 resolution at regions, where there are many stations and hence strong contribution from
165 emissions, while it lowers resolution at regions that lack measurement stations following a
166 method proposed by Stohl et al. (2010)

167 Theoretically, the algorithm can calculate negative posterior emissions, which are
168 physically unlikely. To tackle this problem, an inequality constraint was applied on the
169 emissions following the method of Thacker (2007) that applies the constraint as “error-free”
170 observations:

$$171 \quad \hat{\mathbf{x}} = \mathbf{x} + \mathbf{A}\mathbf{P}^T(\mathbf{P}\mathbf{A}\mathbf{P}^T)^{-1}(\mathbf{c} - \mathbf{P}\mathbf{x}) \quad (4)$$

172 where \mathbf{A} is the posterior error covariance matrix, \mathbf{P} is a matrix operator to select the variables
173 that violate the inequality constraint, and \mathbf{c} is a vector of the inequality constraint, which in this
174 case is zero.

175 We evaluated the assumptions made on the error covariance matrices for the prior
176 emissions and the observations using the reduced χ^2 statistics (\mathbf{B} and \mathbf{R}). When χ^2 is equal to
177 unity, the posterior solution is within the limits of the prescribed uncertainties. The latter is the
178 value of the cost function at the optimum and is equivalent to the weighted sum of squares
179 divided by the number of observations (Thompson et al., 2015). In the inversions performed
180 here, the calculated χ^2 values were between 0.8 and 1.5 indicating that the chosen uncertainty
181 parameters are close to the ideal ones. The number of measurements used in each inversion was
182 equal to 12538 from 17 stations. To select the inversion that provides the most statistically
183 significant result, an evaluation of the improvement in the posterior modelled concentrations,
184 with respect to the prior ones, against the observations was performed (Figure 2). The resulting
185 values of each of the statistical measures that were performed are given in detail in Table 2.
186 Note that this is not a validation of the posterior emissions, because the comparison is only done
187 for the observations that were included in the inversion (dependent observations), and the
188 inversion algorithm has been designed to reduce the model–observation mismatches. This
189 means that the reduction of the posterior concentration mismatches to the observations is
190 determined by the weighting that is given to the observations with respect to the prior emissions.
191 A proper validation of the posterior emissions is performed against observations that were not
192 included in the inversion (independent observations) in Results section.



193 **2.4 Prior emissions**

194 As a priori emissions in the inversions, the ECLIPSE version 5 and 6 (Evaluating the
195 CLimate and Air Quality ImPacts of ShortlivEd Pollutants) (Klimont et al., 2017), EDGAR
196 (Emissions Database for Global Atmospheric Research) version HTAP_v2.2 (Janssens-
197 Maenhout et al., 2015), ACCMIP (Emissions for Atmospheric Chemistry and Climate Model
198 Intercomparison Project) version 5^{ref} (Lamarque et al., 2013) and PKU (Peking University)
199 (Wang et al., 2014b) were used (Figure 3). All inventories include the basic emission sectors
200 (e.g., waste burning, industrial combustion and processing, all means of transportation (aerial,
201 surface, ocean), energy conversion, residential and commercial combustion (see references
202 therein). Biomass burning emissions were adopted from the Global Fire Emissions Database,
203 Version 4.1s (GFEDv4.1s)(Giglio et al., 2013).

204 **2.5 MERRA-2 (Modern-Era Retrospective Analysis for Research and Applications 205 Version 2)**

206 The MERRA-2 reanalysis dataset for BC (Randles et al., 2017) assimilates bias-corrected
207 AOD from Moderate Resolution Imaging Spectroradiometer (MODIS), Advanced Very High
208 Resolution Radiometer (AVHRR) instruments, Multiangle Imaging SpectroRadiometer
209 (MISR) and Aerosol Robotic Network (AERONET) with the Goddard Earth Observing System
210 Model Version 5 (GEOS-5). BC and other aerosols in MERRA-2 are simulated with the
211 Goddard Chemistry, Aerosol, Radiation and Transport (GOCART) model and delivered in
212 hourly to monthly temporal resolution and $0.5^\circ \times 0.625^\circ$ spatial. The product has been validated
213 for AOD, PM and BC extensively (Buchard et al., 2017; Qin et al., 2019; Randles et al., 2017;
214 Sun et al., 2019). Ångström exponent (AE), a measure of how the AOD changes relative to the
215 various wavelength of light, is derived here from AOD₄₆₉, AOD₅₅₀, AOD₆₇₀, and AOD₈₆₅,
216 by fitting the data to the linear transform of Ångström's empirical expression:

$$217 \quad \tau_\lambda = \tau_{\lambda_0} \left(\frac{\lambda}{\lambda_0}\right)^{-\alpha} \quad (5)$$

218 where τ_λ is the known AOD at wavelength λ (in nm), τ_{λ_0} is the AOD at 1000 nm, and α stands
219 for AE (Gueymard and Yang, 2020).

220 **2.6 Absorption Ångström exponent from AErosol RObotic NETWORK (AERONET) 221 data**

222 Aerosol composition over Europe during the COVID-19 lockdown was confirmed using
223 the AERONET data (Holben et al., 1998). AERONET provides globally distributed



224 observations of spectral aerosol optical depth (AOD), inversion products, and precipitable
225 water in diverse aerosol regimes. The AE for a spectral dependence of 440-870 nm is related to
226 the aerosol particle size. Values less than 1 suggest an optical dominance of coarse particles
227 corresponding to dust, ash and sea spray aerosols, while values greater than one imply
228 dominance of fine particles such as smoke and industrial pollution (Eck et al., 1999). We chose
229 data from five stations covering Western, Central and Eastern Europe, for which cloud-free
230 measurements exist for the lockdown period, namely Ben Salem (9.91°E, 35.55°N), Minsk
231 (27.60°E, 53.92°N), Montsec (0.73°E, 42.05°N), MetObs Lindenberg (14.12°E, 52.21°N) and
232 Munich University (11.57°E, 48.15°N). We used Level 1.5 absorption AE (AAE)
233 measurements for the COVID-19 lockdown period (14 March to 30 April 2020).

234 2.7 Statistical measures

235 For the performance evaluation of the inversion results against dependent (observations
236 that were included in the inversion) and independent observations (observations that were not
237 included in the inversion), four different statistical quantities were used:

238 (1) Pearson's correlation coefficient:

$$239 R_{mo} = \frac{n \sum_{i=1}^n m_i o_i - \sum_{i=1}^n m_i \sum_{i=1}^n o_i}{\sqrt{n \sum_{i=1}^n m_i^2 - (\sum_{i=1}^n m_i)^2} \sqrt{n \sum_{i=1}^n o_i^2 - (\sum_{i=1}^n o_i)^2}} \quad (6)$$

240 where n is sample size, m and o the individual sample points for model concentrations and
241 observations indexed with i .

242 (2) The normalized root mean square error (nRMSE):

$$243 nRMSE = \frac{\sqrt{\sum_{i=1}^n \frac{1}{n} (m_i - o_i)^2}}{o_i^{max} - o_i^{min}} \quad (7)$$

244 (3) The mean fractional bias MFB was selected as a symmetric performance indicator that gives
245 equal weights to under- or over-estimated concentrations (minimum to maximum values range
246 from -200% to 200%) and is defined as:

$$247 MFB = \frac{1}{n} \frac{\sum_{i=1}^n (m_i - o_i)}{\sum_{i=1}^n \frac{m_i + o_i}{2}} \quad (8)$$

248 (4) The mean absolute error was computed normalized ($nMAE$) over the average of all the
249 actual values (observations here), which is a widely used simple measure of error:

$$250 nMAE = \frac{\sum_{i=1}^n |m_i - o_i|}{\sum_{i=1}^n o_i} \quad (9)$$



251 **2.8 Region definitions**

252 All country and regional masks are publicly available. Regions used for statistical
253 processing purposes were adopted from the United Nations Statistics Division
254 (<https://unstats.un.org/home/>). Accordingly, Northern Europe includes UK, Norway, Denmark,
255 Sweden, Finland, Iceland, Estonia, Latvia and Lithuania. Southern Europe includes Spain, Italy,
256 Greece, Slovenia, Croatia, Bosnia, Serbia, Albania and North Macedonia. Western Europe is
257 defined by France, Belgium, Holland, Germany, Austria and Switzerland. Eastern Europe
258 includes Poland, Czechia, Slovakia, Hungary, Romania, Bulgaria, Moldova, Ukraine, Belarus
259 and Russia.

260 **3 Results**

261 **3.1 Optimized (posterior) emissions from Bayesian inversion**

262 We performed five inversions for BC over Europe for 1st January- 30th April 2020, each
263 with different prior emissions from ECLIPSE version 5 and 6, EDGAR version HTAP_v2.2,
264 ACCMIP version 5 and PKU (Figure 3). Total prior emissions of BC in Europe from the five
265 emission inventories for the period of the inversion ranged between 192-377 kt. We evaluated
266 the assumptions made on the error covariance matrices for the prior emissions and the
267 observations using the reduced χ^2 statistic (**B** and **R**, see section 2.3). When χ^2 is equal to
268 unity, the posterior solution is within the limits of the prescribed uncertainties. The performance
269 of the inversions with the five different prior inventories was evaluated using four statistical
270 parameters (see section 2.7). The best performance of the inversions was achieved using
271 ECLIPSEv6 (Table 2 and Figure 2) with the smallest *nRMSE* (0.073) value, the largest
272 Pearson's R^2 (0.60), the closest to zero *MFB* value (0.03) and the smallest *nMAE* (714).
273 Therefore, all the results presented below correspond to this inversion.

274 Emissions were calculated to be 191 kt in the inversion domain or approximately 20%
275 smaller than those in ECLIPSEv6 (239 kt) (Figure 4). Note that the aforementioned numbers
276 refer to the whole inversion domain (not only Europe) and the whole study period (January –
277 April 2020). The largest posterior decreases were found in Eastern Europe, where emissions
278 dropped from 35 to 29 kt. Western European emissions of BC declined by almost 11% (from
279 45 to 40 kt), as those in Scandinavian countries (from 8.7 to 6.4 kt). Finally, in Southern Europe
280 (Spain, Italy, Greece) the posterior emissions also decreased by 21% relative to the priors (from
281 61 to 48 kt). The largest country decreases were seen in France (from 14 to 8.2 kt), Italy (from



282 8.0 to 5.9 kt), UK (from 4.4 to 3.1 kt) and Germany (from 4.5 to 4.1 kt). Surprisingly, BC
283 emissions were slightly enhanced in Poland (from 21 to 23 kt), and Spain (from 6.3 to 7.5 kt)
284 were slightly enhanced. Note that although posterior emissions may only imply error in the
285 prior emissions, we demonstrate that the decrease was in fact due to the COVID-19 lockdowns
286 in the next section (3.2).

287 **3.2 Comparison with previous years**

288 We also performed inversions for 2015–2019 for the same period as the 2020 lockdowns
289 (January- April) using almost the same measurement stations and keeping the same settings.
290 The difference in BC emissions during the lockdown in 2020 (14 March to 30 April) from the
291 respective emissions during the same period in 2015–2019 (14 March to 30 April) are shown
292 in Figure 5 (a, emission anomaly) together with the gross domestic product (GDP) (Kummu et
293 al., 2020) (b), and (c) temperature anomaly from ERA-5 (Copernicus Climate Change Service
294 (C3S), 2020) for the same period as the emission anomaly. The difference in the 2020 emissions
295 of BC during the lockdown from the respective emissions in the same period in each of the
296 previous years (2015–2019) is illustrated in Supplementary Figure 1.

297 Overall, BC emissions decreased by ~48 kt during the COVID-19 lockdown in the
298 inversion domain. We record a significant decrease in BC emissions in Central Europe
299 (Northern Italy, Austria, Germany, Spain and some Balkan countries) (Figure 5) compared to
300 the same period during the last five years. On average, emissions declined from 44 to 39 kt over
301 Europe. The decrease has the same characteristics when compared to each of previous years
302 since 2015 (Supplementary Figure 1) based on measurements of BC at the same stations as
303 those used for the 2020 inversion. The countries that showed drastic reductions in BC emissions
304 were those that suffered from the pandemic dramatically, with many human losses, strict social
305 distancing rules and consequently less transport. Specifically, emissions dropped by 20% in
306 Italy (2.4 to 2.0 kt), 40% in Germany (2.2 to 1.3 kt), 23% in Spain (3.4 to 2.6 kt) and remained
307 the same or were slightly enhanced in Poland (~7.0 kt), France (~2.2 kt) and Scandinavia (~1.2
308 kt). Overall, BC emissions in Western Europe declined by 22% (6.2 to 4.8 kt), in Southern
309 Europe by 36% (12 to 7.7 kt) and in Northern Europe by 26% (3.8 to 2.8 kt). BC emissions in
310 Eastern Europe were slightly increased during the 2020 lockdown as compared to the same
311 period in the last five years (22 to 24 kt). Note that these numbers correspond to BC emissions
312 during the COVID-19 lockdown period only (mid-March – April 2020).

313 Some localised areas of increased BC emissions exist in Southern France, Belgium,
314 Northern Germany and Eastern Europe (Figure 5), which are observed relative to almost every



315 year since 2015 (Supplementary Figure 1). While some hotspots in France cannot be easily
316 explained, increased emissions in Eastern European countries are likely due to increased
317 residential combustion, as people had to stay home during the lockdown. The combination of
318 the financial consequences of the COVID-19 lockdown with the relatively low GDP per capita
319 in these countries and the fact that from mid-March to end of April 2020 surface temperatures
320 in these countries were significantly lower than in previous years is suggestive of increased
321 emissions due to residential combustion. This source is most important in Eastern Europe
322 (Klimont et al., 2017). Although residential combustion can be performed for heating or
323 cooking needs in poorer countries, it is also believed to provide a more natural type of warmth
324 and a comfortable and relaxing environment. Hence, it should not be assumed as an emission
325 source in countries with lower GDPs only, especially as people spent more time at home.
326 Moreover, the prevailing average temperatures over Europe during the lockdown were below
327 15°C (Supplementary Figure 2), a temperature used as a basis temperature below which
328 residential combustion increases (Quayle and Diaz, 1980; Stohl et al., 2013).

329 **3.3 Uncertainty and validation of the posterior emissions**

330 One of the basic problems when dealing with inverse modelling is that changing model,
331 observational, or prior uncertainties can have drastic impacts on posterior emissions. We
332 addressed this issue by finding the optimal parameters, in order to have a reduced χ^2 statistic
333 around unity (see section 2.3). However, there are two other sources of uncertainty that,
334 although not linked with the inversion algorithm, could affect posterior emissions drastically.
335 The first is the use of different prior emissions; to estimate this type of uncertainty, we
336 performed five inversions using each of the prior emission datasets (ECLIPSEv6 and v5,
337 EDGAR_HTAPv2.2, ACCMIPv5 and PKU). The uncertainty was calculated as the gridded
338 standard deviation of the posterior emissions resulting from the five inversions. The second
339 type of uncertainty concerns measurement of BC, which is defined as a function of five
340 properties (Petzold et al., 2013). However, as of today, no single instrument exists that could
341 measure all of these properties at the same time. Hence, BC is not a single particle constituent,
342 rather an operational definition depending on the measurement technique (Petzold et al., 2013).
343 Here we use light absorption coefficients (Petzold et al., 2013) converted to equivalent BC
344 (eBC) using the mass absorption cross section (MAC). The MAC is instrument specific and
345 wavelength dependent. The site-specific MAC values used to convert the filter-based light
346 absorption to eBC can be seen in Table 1. It has been reported that MAC values vary from 2 –
347 3 m² g⁻¹ up to 20 m² g⁻¹ (Bond and Bergstrom, 2006). To estimate the uncertainty of the posterior



348 fluxes associated with the variable MAC, we performed a sensitivity study using MAC values
349 of 5, 10 and 20 $\text{m}^2 \text{g}^{-1}$ in all stations, as well as variable MAC values for each stations (Table
350 1). Since these values are lognormally distributed, the uncertainty is calculated as the geometric
351 standard deviation. The impact of other sources of uncertainty, such as those referring to
352 scavenging coefficients, particle size and density that are used in the model have been studied
353 before (Evangelidou et al., 2018; Grythe et al., 2017).

354 The posterior emissions are less sensitive to the use of different MACs than the use of
355 different prior inventories (Figure 6). The relative uncertainty due to different use of MAC
356 values was up to 20–30% in most of Europe and increases dramatically far from the
357 observations. Accordingly, the emission uncertainty of BC from the use of different priors was
358 estimated to be up to 40% in Europe and shows very similar characteristics (same hot-spot
359 regions and larger values where measurements lack). Overall, the combined uncertainty of BC
360 emissions was ~60% in Europe.

361 Validation of top-down emissions obtained by inversion algorithms can be proper only if
362 measurements that were not included in the inversion are to be used (independent observations).
363 For this reason, we left out of the inversion observations from two stations (DE0054K and
364 DE0066R, Table 1). Due to the higher measurement station density in Central Europe, we
365 randomly selected two German stations, rather than from a country that is adjacent to regions
366 that lack observations.

367 The prior, optimized and measured concentrations are shown in Figure 7 together with
368 MERRA-2 surface BC concentrations at the same stations. The average footprint emission
369 sensitivities are also given for the period of the lockdown. At station DE0054K, prior emissions
370 represent observations very well until the beginning of the lockdown and then fail (Figure 7).
371 On the other hand, the posterior emissions represent the variant concentrations during the
372 lockdown effectively and also manage to capture some concentration peaks, which is reflected
373 by lower *nRMSE*. Backward modelling showed that the enhanced concentrations originate
374 from Northern Germany and the Netherlands, where posterior emissions were increased
375 compared with the prior ones (Figure 4). A similar pattern was seen at station DE0066K,
376 although this station showed concentrations up to 4 mg m^{-3} (Figure 7). Again, the optimized
377 emissions managed to represent the peaks at the end of January 2020 and at the beginning of
378 the lockdown, which is again reflected by a half *nRMSE* values and *MFB* close to zero as
379 compared to the priors. The larger concentrations during the lockdown result from increased
380 emissions over Eastern Germany, Poland and the Netherlands, as well as in oil industries in the



381 North Sea (Figure 4b). In all these regions the footprint emissions sensitivities corresponding
382 to the two independent stations were the highest.

383 4 Discussion

384 The improved air quality that Europe experienced during the lockdown was also evident
385 from the assimilated MERRA-2 satellite-based BC data. The latter are plotted in Supplementary
386 Figure 3 (left axis), together with the posterior emissions calculated in the present study (right
387 axis) between 2015–2020. Many of the ACTRIS stations reported increased light absorption in
388 the beginning of the lockdown (e.g., Figure 7); MERRA-2 data show the same patterns in
389 France, Italy, UK and in Spain, and in all of Europe, in general. This can be explained by
390 residential combustion considering that the surface temperature during the lockdown was lower
391 than in previous years (Figure 5). The latter was confirmed by MERRA-2 reanalysis Ångström
392 Exponent (AE) parameter at 470–870 nm, which shows higher values over Central and Eastern
393 Europe during the lockdown in 2020 than in the same period of the previous years (Figure 8a,b).
394 Larger AE values confirm the presence of wood burning aerosols (Eck et al., 1999). The fact
395 that during the COVID-19 lockdown, residential combustion was a significant aerosol source
396 in Europe, as compared to the previous years, was also confirmed by real-time observations of
397 absorption AE from the AERONET data in five selected stations over Europe (Figure 8c).
398 Measured absorption AE was higher during mid-March to April 2020 than in the same period
399 of the last five years.

400 Emissions of BC calculated with Bayesian inversion for the lockdown period dropped
401 substantially in most of the countries that suffered from further spread of the virus and,
402 accordingly, from strict lockdown measures, as compared to the respective emissions in the
403 beginning of 2020 (Supplementary Figure 3). Specifically, the decrease in France was as high
404 as 40%, 11% in Italy, 32% in Germany, 15% in Spain and 25% in the UK. Emissions also
405 declined in Scandinavia by 12%, although Sweden did not enforce a lockdown. Overall, a
406 reduction in BC emissions of about 10% can be concluded for Europe as a whole due to the
407 lockdown. Stronger decreases in Eastern Europe were likely partly compensated by increased
408 residential combustion in resulting from the prevailing low temperatures.

409 We report a decrease of 11% in BC emissions in Europe during the lockdown that resulted
410 from the COVID-19 outbreak, as compared to the same period in all previous years since 2015,
411 based on particle light absorption measurements. We highlight these changes in BC emissions
412 as a result of COVID-19 restrictions by plotting the temporal variability of the BC emissions



413 in the 5 previous years (2015 – 2019) for France, Italy, Germany, Spain, Scandinavia and
414 Europe (Figure 9). We record a decrease in emissions of BC in France, Italy, Germany and
415 Scandinavia in mid-March to April 2020, opposite to what was estimated for all years between
416 2015 and 2019, which is obviously due to COVID-19. The UK and Spain showed a similar
417 decrease in mid-March to April emissions in all previous years (2015–2019). However, the
418 estimated posterior BC emissions during the 2020 lockdown were significantly lower than those
419 of the same period in any of the previous years. Overall, emissions declined by 20% in Italy,
420 40% in Germany, 23% in Spain and remained the same in France and slightly enhanced in
421 Scandinavia as compared to those of the last five years.

422 5 Conclusions

423 The impact of the COVID-19 lockdowns over Europe on the BC emissions, in response
424 to the pandemic was assessed in the present manuscript. Particle light absorption measurements
425 from 17 ACTRIS stations all around Europe were rapidly gathered and cleaned to produce a
426 high-quality product. The latter was used in a well-established Bayesian inversion framework
427 and BC emissions were optimised over Europe to better capture the observations. We calculate
428 that the optimised (posterior) BC emissions declined from 44 to 39 kt (11%) during the
429 lockdown over Europe, as compared to the same period in the previous five years (2015–2019).
430 The largest reductions were calculated for countries that suffered from the pandemic
431 dramatically, such as Italy (from 2.4 to 2.0 kt), Germany (from 2.2 to 1.3 kt), Spain (from 3.4
432 to 2.6 kt). BC emissions in Western Europe during the 2020 lockdown were decreased from 6.2
433 to 4.8 kt (22%), in Southern Europe from 12 to 7.7 kt (36%) and in Northern Europe from 3.8
434 to 2.8 kt (26%) as compared to the same period in the last five years. BC emissions were slightly
435 enhanced in Eastern Europe (from 22 to 24 kt) during the lockdown, due to increased residential
436 combustion, as people had to stay home and temperatures at that time were the lowest of the
437 last five years. The presence of wood burning aerosols during the lockdowns was confirmed by
438 large MERRA-2 AE, as well as from absorption AE measurements from AERONET that were
439 higher in the lockdowns than in the same period of the last five years. The impact of the
440 European lockdowns on BC emissions was also confirmed by a 10% decrease of the posterior
441 emissions over Europe in the beginning of the lockdown, as compared to the period before,
442 opposite to what was calculated in the previous years, which is obviously due to COVID-19.
443 This decrease was more pronounced in France (40%), Italy (11%), Germany (32%), Spain
444 (15%), UK (25%) and in Scandinavian countries (12%). The full impact of the disastrous
445 pandemic will likely take years to assess. Nevertheless, with COVID-19 cases once again



446 increasing in many countries, the information presented here are essential to understand the full
447 health and climate impacts of lockdown measures.

448 *Data availability.* All measurement data and model outputs used for the present publication are
449 open and can be downloaded from <https://doi.org/10.21336/gen.b5vj-sn33> or upon request to
450 the corresponding author. All prior emission datasets are also available for download. ECLIPSE
451 emissions can be obtained from
452 http://www.iiasa.ac.at/web/home/research/researchPrograms/air/Global_emissions.html,
453 EDGAR version HTAP_V2.2 from <http://edgar.jrc.ec.europa.eu/methodology.php#>, ACCMIP
454 version 5 from http://accent.aero.jussieu.fr/ACCMIP_metadata.php) and PKU from
455 <http://inventory.pku.edu.cn>. FLEXPART is open in public and can be downloaded from
456 <https://www.flexpart.eu>, so as FLEXINVERT+ from <https://flexinvert.nilu.no>. MERRA-2 re-
457 analysis data can be obtained from <https://disc.gsfc.nasa.gov>, so as AERONET measurements
458 from <https://aeronet.gsfc.nasa.gov>.

459 *Author contributions.* N.E. led the work and wrote the paper. S.E. and A.S. commented on the
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479

480 References

- 481 Ackerman, a. S.: Reduction of Tropical Cloudiness by Soot, *Science* (80-.), 288(5468),
482 1042–1047, doi:10.1126/science.288.5468.1042, 2000.
- 483 Adams, M. D.: Air pollution in Ontario, Canada during the COVID-19 State of Emergency,
484 *Sci. Total Environ.*, 742, 140516, doi:10.1016/j.scitotenv.2020.140516, 2020.
- 485 Bauwens, M., Compernelle, S., Stavrakou, T., Müller, J. F., van Gent, J., Eskes, H., Levelt, P.
486 F., van der A, R., Veeffkind, J. P., Vlietinck, J., Yu, H. and Zehner, C.: Impact of Coronavirus
487 Outbreak on NO₂ Pollution Assessed Using TROPOMI and OMI Observations, *Geophys.*
488 *Res. Lett.*, 47(11), 1–9, doi:10.1029/2020GL087978, 2020.
- 489 Berman, J. D. and Ebisu, K.: Changes in U.S. air pollution during the COVID-19 pandemic,
490 *Sci. Total Environ.*, 739, 139864, doi:10.1016/j.scitotenv.2020.139864, 2020.
- 491 Bond, T. C. and Bergstrom, R. W.: Light Absorption by Carbonaceous Particles: An
492 Investigative Review, *Aerosol Sci. Technol.*, 40(1), 27–67, doi:10.1080/02786820500421521,
493 2006.
- 494 Bond, T. C., Doherty, S. J., Fahey, D. W., Forster, P. M., Berntsen, T., Deangelo, B. J.,
495 Flanner, M. G., Ghan, S., Kärcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P. K., Sarofim,
496 M. C., Schultz, M. G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S., Bellouin, N.,
497 Guttikunda, S. K., Hopke, P. K., Jacobson, M. Z., Kaiser, J. W., Klimont, Z., Lohmann, U.,
498 Schwarz, J. P., Shindell, D., Storelvmo, T., Warren, S. G. and Zender, C. S.: Bounding the
499 role of black carbon in the climate system: A scientific assessment, *J. Geophys. Res. Atmos.*,
500 118(11), 5380–5552, doi:10.1002/jgrd.50171, 2013.
- 501 Buchard, V., Randles, C. A., da Silva, A. M., Darmenov, A., Colarco, P. R., Govindaraju, R.,
502 Ferrare, R., Hair, J., Beyersdorf, A. J., Ziemba, L. D. and Yu, H.: The MERRA-2 aerosol
503 reanalysis, 1980 onward. Part II: Evaluation and case studies, *J. Clim.*, 30(17), 6851–6872,
504 doi:10.1175/JCLI-D-16-0613.1, 2017.
- 505 Cassiani, M., Stohl, A. and Brioude, J.: Lagrangian Stochastic Modelling of Dispersion in the
506 Convective Boundary Layer with Skewed Turbulence Conditions and a Vertical Density
507 Gradient: Formulation and Implementation in the FLEXPART Model, *Boundary-Layer*
508 *Meteorol.*, 154(3), 367–390, doi:10.1007/s10546-014-9976-5, 2014.
- 509 Clarke, A. D. and Noone, K. J.: Soot in the arctic snowpack: a cause for perturbations in
510 radiative transfer, *Atmos. Environ.*, 41(SUPPL.), 64–72, doi:10.1016/0004-6981(85)90113-1,
511 1985.
- 512 Conticini, E., Frediani, B. and Caro, D.: Can atmospheric pollution be considered a co-factor
513 in extremely high level of SARS-CoV-2 lethality in Northern Italy?, *Environ. Pollut.*, 261,
514 114465, doi:10.1016/j.envpol.2020.114465, 2020.
- 515 Copernicus Climate Change Service (C3S): C3S ERA5-Land reanalysis . Copernicus Climate
516 Change Service, [online] Available from: <https://cds.climate.copernicus.eu/cdsapp#!/home>
517 (Accessed 31 August 2020), 2020.
- 518 Dantas, G., Siciliano, B., França, B. B., da Silva, C. M. and Arbilla, G.: The impact of
519 COVID-19 partial lockdown on the air quality of the city of Rio de Janeiro, Brazil, *Sci. Total*
520 *Environ.*, 729, doi:10.1016/j.scitotenv.2020.139085, 2020.
- 521 Diffenbaugh, N. S., Field, C. B., Appel, E. A., Azevedo, I. L., Baldocchi, D. D., Burke, M.,
522 Burney, J. A., Ciais, P., Davis, S. J., Fiore, A. M., Fletcher, S. M., Hertel, T. W., Horton, D.
523 E., Hsiang, S. M., Jackson, R. B., Jin, X., Levi, M., Lobell, D. B., McKinley, G. A., Moore, F.



- 524 C., Montgomery, A., Nadeau, K. C., Pataki, D. E., Randerson, J. T., Reichstein, M., Schnell,
525 J. L., Seneviratne, S. I., Singh, D., Steiner, A. L. and Wong-Parodi, G.: The COVID-19
526 lockdowns: a window into the Earth System, *Nat. Rev. Earth Environ.*, 1–12,
527 doi:10.1038/s43017-020-0079-1, 2020.
- 528 Dutheil, F., Baker, J. S. and Navel, V.: COVID-19 as a factor influencing air pollution?,
529 *Environ. Pollut.*, 263, 2019–2021, doi:10.1016/j.envpol.2020.114466, 2020.
- 530 Eck, T. F., Holben, B. N., Reid, J. S., Smirnov, A., Neill, N. T. O., Slutsker, I. and Kinne, S.:
531 Wavelength dependence of the optical depth of biomass burning, urban, and desert dust
532 aerosols, *J. Geophys. Res.*, 104(D24), 31333–31349, 1999.
- 533 Evangeliou, N., Thompson, R. L., Eckhardt, S. and Stohl, A.: Top-down estimates of black
534 carbon emissions at high latitudes using an atmospheric transport model and a Bayesian
535 inversion framework, *Atmos. Chem. Phys.*, 18(20), doi:10.5194/acp-18-15307-2018, 2018.
- 536 Forster, C., Stohl, A. and Seibert, P.: Parameterization of convective transport in a Lagrangian
537 particle dispersion model and its evaluation, *J. Appl. Meteorol. Climatol.*, 46(4), 403–422,
538 doi:10.1175/JAM2470.1, 2007.
- 539 Giglio, L., Randerson, J. T. and van der Werf, G. R.: Analysis of daily, monthly, and annual
540 burned area using the fourth-generation global fire emissions database (GFED4), *J. Geophys.*
541 *Res. Biogeosciences*, 118, 317–328, doi:10.1002/jgrg.20042, 2013, 2013.
- 542 Grythe, H., Kristiansen, N. I., Groot Zwaaftink, C. D., Eckhardt, S., Ström, J., Tunved, P.,
543 Krejci, R. and Stohl, A.: A new aerosol wet removal scheme for the Lagrangian particle
544 model FLEXPARTv10, *Geosci. Model Dev.*, 10, 1447–1466, doi:10.5194/gmd-10-1447-
545 2017, 2017.
- 546 Gueymard, C. A. and Yang, D.: Worldwide validation of CAMS and MERRA-2 reanalysis
547 aerosol optical depth products using 15 years of AERONET observations, *Atmos. Environ.*,
548 225(November 2019), 117216, doi:10.1016/j.atmosenv.2019.117216, 2020.
- 549 He, G., Pan, Y. and Tanaka, T.: The short-term impacts of COVID-19 lockdown on urban air
550 pollution in China, *Nat. Sustain.*, doi:10.1038/s41893-020-0581-y, 2020.
- 551 Hegg, D. A., Warren, S. G., Grenfell, T. C., Doherty, S. J., Larson, T. V. and Clarke, A. D.:
552 Source attribution of black carbon in arctic snow, *Environ. Sci. Technol.*, 43(11), 4016–4021,
553 doi:10.1021/es803623f, 2009.
- 554 Holben, B. N., Eck, T. F., Slutsker, I., Tanré, D., Buis, J. P., Setzer, A., Vermote, E., Reagan,
555 J. A., Kaufman, Y. J., Nakajima, T., Lavenu, F., Jankowiak, I. and Smirnov, A.:
556 AERONET—A Federated Instrument Network and Data Archive for Aerosol
557 Characterization, *Remote Sens. Environ.*, 66(1), 1–16, doi:10.1016/S0034-4257(98)00031-5,
558 1998.
- 559 Huang, C., Wang, Y., Li, X., Ren, L., Zhao, J., Hu, Y., Zhang, L., Fan, G., Xu, J., Gu, X.,
560 Cheng, Z., Yu, T., Xia, J., Wei, Y., Wu, W., Xie, X., Yin, W., Li, H., Liu, M., Xiao, Y., Gao,
561 H., Guo, L., Xie, J., Wang, G., Jiang, R., Gao, Z., Jin, Q., Wang, J. and Cao, B.: Clinical
562 features of patients infected with 2019 novel coronavirus in Wuhan, China, *Lancet*,
563 395(10223), 497–506, doi:10.1016/S0140-6736(20)30183-5, 2020.
- 564 Janssens-Maenhout, G., Crippa, M., Guizzardi, D., Dentener, F., Muntean, M., Pouliot, G.,
565 Keating, T., Zhang, Q., Kurokawa, J., Wankmüller, R., Denier Van Der Gon, H., Kuenen, J. J.
566 P., Klimont, Z., Frost, G., Darras, S., Koffi, B. and Li, M.: HTAP-v2.2: A mosaic of regional
567 and global emission grid maps for 2008 and 2010 to study hemispheric transport of air
568 pollution, *Atmos. Chem. Phys.*, 15(19), 11411–11432, doi:10.5194/acp-15-11411-2015, 2015.
- 569 Jinhuan, Q. and Liqian, Y.: Variation characteristics of atmospheric aerosol optical depths
570 and visibility in North China during 1980 } 1994, *Atmos. Environ.*, 34, 603–609, 2000.
- 571 John Hopkins University of Medicine: Coronavirus resource center, [online] Available from:
572 <https://coronavirus.jhu.edu/map.html> (Accessed 10 August 2020), 2020.
- 573 Kerimray, A., Baimatova, N., Ibragimova, O. P., Bukenov, B., Kenessov, B., Plotitsyn, P. and



- 574 Karaca, F.: Assessing air quality changes in large cities during COVID-19 lockdowns: The
575 impacts of traffic-free urban conditions in Almaty, Kazakhstan, *Sci. Total Environ.*, 730,
576 139179, doi:10.1016/j.scitotenv.2020.139179, 2020.
- 577 Klimont, Z., Kupiainen, K., Heyes, C., Purohit, P., Cofala, J., Rafaj, P., Borcken-Kleefeld, J.
578 and Schöpp, W.: Global anthropogenic emissions of particulate matter including black
579 carbon, *Atmos. Chem. Phys.*, 17, 8681–8723, doi:10.5194/acp-17-508681-2017, 2017.
- 580 Kummu, M., Taka, M. and Guillaume, J. H. A.: Data from: Gridded global datasets for Gross
581 Domestic Product and Human Development Index over 1990-2015, v2, Dryad, Dataset,
582 doi:https://doi.org/10.5061/dryad.dk1j0, 2020.
- 583 Lack, D. A. and Langridge, J. M.: On the attribution of black and brown carbon light
584 absorption using the Ångström exponent, *Atmos. Chem. Phys.*, 13(20), 10535–10543,
585 doi:10.5194/acp-13-10535-2013, 2013.
- 586 Laj, P., Bigi, A., Rose, C., Andrews, E., Lund Myhre, C., Collaud Coen, M., Wiedensohler,
587 A., Schultz, M., Ogren, J., Fiebig, M., Glib, J., Mortier, A., Pandolfi, M., Petäjä, T., Kim, S.-
588 W., Aas, W., Putaud, J.-P., Mayol-Bracero, O., Keywood, M., Labrador, L., Aalto, P.,
589 Ahlberg, E., Alados Arboledas, L., Alastuey, A., Andrade, M., Artñano, B., Ausmeel, S.,
590 Arsov, T., Asmi, E., Backman, J., Baltensperger, U., Bastian, S., Bath, O., Beukes, J. P.,
591 Brem, B., Bukowiecki, N., Conil, S., Couret, C., Day, D., Dayantolis, W., Degorska, A., Dos
592 Santos, S. M., Eleftheriadis, K., Fetfatzis, P., Favez, O., Flentje, H., Gini, M., Gregorič, A.,
593 Gysel-Beer, M., Hallar, G., Hand, J., Hoffer, A., Hueglin, C., Hooda, R., Hyvärinen, A.,
594 Kalapov, I., Kalivitis, N., Kasper-Giebl, A., Kim, J. E., Kouvarakis, G., Kranjc, I., Krejci, R.,
595 Kulmala, M., Labuschagne, C., Lee, H.-J., Lihavainen, H., Lin, N.-H., Löschau, G., Luoma,
596 K., Marinoni, A., Meinhardt, F., Merkel, M., Metzger, J.-M., Mihalopoulos, N., Nguyen, N.
597 A., Ondracek, J., Pérez, N., Perrone, M. R., Petit, J.-E., Picard, D., Pichon, J.-M., Pont, V.,
598 Prats, N., Prenni, A., Reisen, F., Romano, S., Sellegri, K., Sharma, S., Schauer, G., Sheridan,
599 P., Sherman, J. P., Schütze, M., Schwerin, A., Sohmer, R., Sorribas, M., Steinbacher, M.,
600 Sun, J., Titos, G., Tokzko, B., et al.: A global analysis of climate-relevant aerosol properties
601 retrieved from the network of GAW near-surface observatories, *Atmos. Meas. Tech. Discuss.*,
602 1–70, doi:10.5194/amt-2019-499, 2020.
- 603 Lamarque, J. F., Shindell, D. T., Josse, B., Young, P. J., Cionni, I., Eyring, V., Bergmann, D.,
604 Cameron-Smith, P., Collins, W. J., Doherty, R., Dalsoren, S., Faluvegi, G., Folberth, G.,
605 Ghan, S. J., Horowitz, L. W., Lee, Y. H., MacKenzie, I. A., Nagashima, T., Naik, V.,
606 Plummer, D., Righi, M., Rumbold, S. T., Schulz, M., Skeie, R. B., Stevenson, D. S., Strode,
607 S., Sudo, K., Szopa, S., Voulgarakis, A. and Zeng, G.: The atmospheric chemistry and climate
608 model intercomparison Project (ACCMIP): Overview and description of models, simulations
609 and climate diagnostics, *Geosci. Model Dev.*, 6(1), 179–206, doi:10.5194/gmd-6-179-2013,
610 2013.
- 611 Le, T., Wang, Y., Liu, L., Yang, J., Yung, Y. L., Li, G. and Seinfeld, J. H.: Unexpected air
612 pollution with marked emission reductions during the COVID-19 outbreak in China, *Science*
613 (80-.), (2), eabb7431, doi:10.1126/science.abb7431, 2020.
- 614 Lian, X., Huang, J., Huang, R., Liu, C., Wang, L. and Zhang, T.: Impact of city lockdown on
615 the air quality of COVID-19-hit of Wuhan city, *Sci. Total Environ.*, 742, 140556,
616 doi:10.1016/j.scitotenv.2020.140556, 2020.
- 617 Long, C. M., Nascarella, M. A. and Valberg, P. A.: Carbon black vs. black carbon and other
618 airborne materials containing elemental carbon: Physical and chemical distinctions, *Environ.*
619 *Pollut.*, 181, 271–286, doi:10.1016/j.envpol.2013.06.009, 2013.
- 620 Myhre, G., Samset, B. H., Schulz, M., Balkanski, Y., Bauer, S., Bernsten, T. K., Bian, H.,
621 Bellouin, N., Chin, M., Diehl, T., Easter, R. C., Feichter, J., Ghan, S. J., Hauglustaine, D.,
622 Iversen, T., Kinne, S., Kirkevåg, A., Lamarque, J. F., Lin, G., Liu, X., Lund, M. T., Luo, G.,
623 Ma, X., Van Noije, T., Penner, J. E., Rasch, P. J., Ruiz, A., Seland, Skeie, R. B., Stier, P.,



- 624 Takemura, T., Tsigaridis, K., Wang, P., Wang, Z., Xu, L., Yu, H., Yu, F., Yoon, J. H., Zhang,
625 K., Zhang, H. and Zhou, C.: Radiative forcing of the direct aerosol effect from AeroCom
626 Phase II simulations, *Atmos. Chem. Phys.*, 13(4), 1853–1877, doi:10.5194/acp-13-1853-2013,
627 2013.
- 628 Otmani, A., Benchrif, A., Tahri, M., Bounakhla, M., Chakir, E. M., El Bouch, M. and
629 Krombi, M.: Impact of Covid-19 lockdown on PM10, SO2 and NO2 concentrations in Salé
630 City (Morocco), *Sci. Total Environ.*, 735(2), 139541, doi:10.1016/j.scitotenv.2020.139541,
631 2020.
- 632 Petzold, A., Ogren, J. A., Fiebig, M., Laj, P., Li, S. M., Baltensperger, U., Holzer-Popp, T.,
633 Kinne, S., Pappalardo, G., Sugimoto, N., Wehrli, C., Wiedensohler, A. and Zhang, X. Y.:
634 Recommendations for reporting black carbon measurements, *Atmos. Chem. Phys.*, 13(16),
635 8365–8379, doi:10.5194/acp-13-8365-2013, 2013.
- 636 Pissou, I., Sollum, E., Grythe, H., Kristiansen, N., Cassiani, M., Eckhardt, S., Arnold, D.,
637 Morton, D., Thompson, R. L., Groot Zwaafink, C. D., Evangeliou, N., Sodemann, H.,
638 Haimberger, L., Henne, S., Brunner, D., Burkhardt, J. F., Fouilloux, A., Brioude, J., Philipp,
639 A., Seibert, P. and Stohl, A.: The Lagrangian particle dispersion model FLEXPART version
640 10.4, *Geosci. Model Dev.*, 12, 4955–4997, doi:10.5194/gmd-12-4955-2019, 2019.
- 641 Qin, W., Zhang, Y., Chen, J., Yu, Q., Cheng, S., Li, W., Liu, X. and Tian, H.: Variation,
642 sources and historical trend of black carbon in Beijing, China based on ground observation
643 and MERRA-2 reanalysis data, *Environ. Pollut.*, 245(2), 853–863,
644 doi:10.1016/j.envpol.2018.11.063, 2019.
- 645 Quayle, R. G. and Diaz, H. F.: Heating degree day data applied to residential heating energy
646 consumption, *J. Appl. Meteorol.*, 19, 241–246, 1980.
- 647 Le Quéré, C., Jackson, R. B., Jones, M. W., Smith, A. J. P., Abernethy, S., Andrew, R. M.,
648 De-Gol, A. J., Willis, D. R., Shan, Y., Canadell, J. G., Friedlingstein, P., Creutzig, F. and
649 Peters, G. P.: Temporary reduction in daily global CO2 emissions during the COVID-19
650 forced confinement, *Nat. Clim. Chang.*, 10(7), 647–653, doi:10.1038/s41558-020-0797-x,
651 2020.
- 652 Randles, C. A., da Silva, A. M., Buchard, V., Colarco, P. R., Darmenov, A., Govindaraju, R.,
653 Smirnov, A., Holben, B., Ferrare, R., Hair, J., Shinozuka, Y. and Flynn, C. J.: The MERRA-2
654 aerosol reanalysis, 1980 onward. Part I: System description and data assimilation evaluation,
655 *J. Clim.*, 30(17), 6823–6850, doi:10.1175/JCLI-D-16-0609.1, 2017.
- 656 Sicard, P., De Marco, A., Agathokleous, E., Feng, Z., Xu, X., Paoletti, E., Rodriguez, J. J. D.
657 and Calatayud, V.: Amplified ozone pollution in cities during the COVID-19 lockdown, *Sci.*
658 *Total Environ.*, 735, doi:10.1016/j.scitotenv.2020.139542, 2020.
- 659 Sohrabi, C., Alsafi, Z., O’Neill, N., Khan, M., Kerwan, A., Al-Jabir, A., Iosifidis, C. and
660 Agha, R.: World Health Organization declares global emergency: A review of the 2019 novel
661 coronavirus (COVID-19), *Int. J. Surg.*, 76(February), 71–76, doi:10.1016/j.ijso.2020.02.034,
662 2020.
- 663 Stohl, A., Forster, C., Frank, A., Seibert, P. and Wotawa, G.: Technical note: The Lagrangian
664 particle dispersion model FLEXPART version 6.2, *Atmos. Chem. Phys.*, 5(9), 2461–2474,
665 doi:10.5194/acp-5-2461-2005, 2005.
- 666 Stohl, A., Kim, J., Li, S., O’Doherty, S., Salameh, P. K., Saito, T., Vollmer, M. K., Wan, D.,
667 Yao, B., Yokouchi, Y. and Zhou, L. X.: Hydrochlorofluorocarbon and hydrofluorocarbon
668 emissions in East Asia determined by inverse modeling, *Atmos. Chem. Phys. Discuss.*, 10(2),
669 2089–2129, doi:10.5194/acpd-10-2089-2010, 2010.
- 670 Stohl, A., Klimont, Z., Eckhardt, S., Kupiainen, K., Shevchenko, V. P., Kopeikin, V. M. and
671 Novigatsky, A. N.: Black carbon in the Arctic: The underestimated role of gas flaring and
672 residential combustion emissions, *Atmos. Chem. Phys.*, 13(17), 8833–8855, doi:10.5194/acp-
673 13-8833-2013, 2013.



- 674 Sun, E., Xu, X., Che, H., Tang, Z., Gui, K., An, L., Lu, C. and Shi, G.: Variation in MERRA-
675 2 aerosol optical depth and absorption aerosol optical depth over China from 1980 to 2017, *J.*
676 *Atmos. Solar-Terrestrial Phys.*, 186(January), 8–19, doi:10.1016/j.jastp.2019.01.019, 2019.
- 677 Tarantola, A.: *Inverse Problem Theory and Methods for Model Parameter Estimation*, Society
678 for Industrial and Applied Mathematics, Philadelphia, Pa., 2005.
- 679 Thacker, W. C.: Data assimilation with inequality constraints, *Ocean Model.*, 16(3–4), 264–
680 276, doi:10.1016/j.ocemod.2006.11.001, 2007.
- 681 Thompson, R. L. and Stohl, A.: FLEXINVERT: An atmospheric Bayesian inversion
682 framework for determining surface fluxes of trace species using an optimized grid, *Geosci.*
683 *Model Dev.*, 7(5), 2223–2242, doi:10.5194/gmd-7-2223-2014, 2014.
- 684 Thompson, R. L., Stohl, A., Zhou, L. X., Dlugokencky, E., Fukuyama, Y., Tohjima, Y., Kim,
685 S. Y., Lee, H., Nisbet, E. G., Fisher, R. E., Lowry, D., Weiss, R. F., Prinn, R. G., O’Doherty,
686 S., Young, D. and White, J. W. C.: Methane emissions in East Asia for 2000–2011 estimated
687 using an atmospheric Bayesian inversion, *J. Geophys. Res. Atmos.*, 120(9), 4352–4369,
688 doi:10.1002/2014JD022394, 2015.
- 689 Wang, P., Wang, H., Wang, Y. Q., Zhang, X. Y., Gong, S. L., Xue, M., Zhou, C. H., Liu, H.
690 L., An, X. Q., Niu, T. and Cheng, Y. L.: Inverse modeling of black carbon emissions over
691 China using ensemble data assimilation, *Atmos. Chem. Phys.*, 16(2), 989–1002,
692 doi:10.5194/acp-16-989-2016, 2016.
- 693 Wang, R., Tao, S., Balkanski, Y., Ciais, P., Boucher, O., Liu, J., Piao, S., Shen, H., Vuolo, M.
694 R., Valari, M., Chen, H., Chen, Y., Cozic, A., Huang, Y., Li, B., Li, W., Shen, G., Wang, B.
695 and Zhang, Y.: Exposure to ambient black carbon derived from a unique inventory and high-
696 resolution model., *Proc. Natl. Acad. Sci. U. S. A.*, 111(7), 2459–63,
697 doi:10.1073/pnas.1318763111, 2014a.
- 698 Wang, R., Tao, S., Shen, H., Huang, Y., Chen, H., Balkanski, Y., Boucher, O., Ciais, P.,
699 Shen, G., Li, W., Zhang, Y., Chen, Y., Lin, N., Su, S., Li, B., Liu, J. and Liu, W.: Trend in
700 global black carbon emissions from 1960 to 2007, *Environ. Sci. Technol.*, 48(12), 6780–6787,
701 doi:10.1021/es5021422, 2014b.
- 702 WHO: Report of the WHO-China Joint Mission on Coronavirus Disease 2019 (COVID-19),
703 WHO-China Jt. Mission Coronavirus Dis. 2019, 2019(February), 16–24 [online] Available
704 from: [https://www.who.int/docs/default-source/coronaviruse/who-china-joint-mission-on-](https://www.who.int/docs/default-source/coronaviruse/who-china-joint-mission-on-covid-19-final-report.pdf)
705 [covid-19-final-report.pdf](https://www.who.int/docs/default-source/coronaviruse/who-china-joint-mission-on-covid-19-final-report.pdf), 2020.
- 706 Zanatta, M., Gysel, M., Bukowiecki, N., Müller, T., Weingartner, E., Areskoug, H., Fiebig,
707 M., Yttri, K. E., Mihalopoulos, N., Kouvarakis, G., Beddows, D., Harrison, R. M., Cavalli, F.,
708 Putaud, J. P., Spindler, G., Wiedensohler, A., Alastuey, A., Pandolfi, M., Sellegri, K.,
709 Swietlicki, E., Jaffrezo, J. L., Baltensperger, U. and Laj, P.: A European aerosol
710 phenomenology-5: Climatology of black carbon optical properties at 9 regional background
711 sites across Europe, *Atmos. Environ.*, 145, 346–364, doi:10.1016/j.atmosenv.2016.09.035,
712 2016.
- 713 Zheng, H., Kong, S., Chen, N., Yan, Y., Liu, D., Zhu, B., Xu, K., Cao, W., Ding, Q., Lan, B.,
714 Zhang, Z., Zheng, M., Fan, Z., Cheng, Y., Zheng, S., Yao, L., Bai, Y., Zhao, T. and Qi, S.:
715 Significant changes in the chemical compositions and sources of PM_{2.5} in Wuhan since the
716 city lockdown as COVID-19, *Sci. Total Environ.*, 739, doi:10.1016/j.scitotenv.2020.140000,
717 2020.
- 718 Zotter, P., Herich, H., Gysel, M., El-Haddad, I., Zhang, Y., Mocnik, G., Hüglin, C.,
719 Baltensperger, U., Szidat, S. and Prévôt, A. S. H.: Evaluation of the absorption Ångström
720 exponents for traffic and wood burning in the Aethalometer-based source apportionment
721 using radiocarbon measurements of ambient aerosol, *Atmos. Chem. Phys.*, 17(6), 4229–4249,
722 doi:10.5194/acp-17-4229-2017, 2017.
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725 **TABLES & FIGURES**

726

727 **Table 1.** Observation sites from the ACTRIS platform used to perform the inversions
728 (dependent observations) and to validate the posterior emissions (independent observations)
729 (the altitude indicates the sampling height in meters above sea level). Multi-Angle Absorption
730 Photometers (MAAP) were used at all sites, except El Arenosillo (ES0100R) where a
731 Continuous Light Absorption Photometer (CLAP) was used, Birkenes (NO0002R), where a
732 Particle Soot Absorption Photometer (PSAP) and Observatoire Perenne de l' Environnement
733 (FR0022R) and Zeppelin (NO0042G) where Aethalometers (AW31) were used.

Name	Latitude	Longitude	Altitude	Type	Wavelength (nm)	MAC _{@637} (m ² g ⁻¹)
Jungfraujoch (CH0001G)	46.55	7.99	3578	Dependent	637	10
Hohenpeissenberg (DE0043G)	47.80	11.01	985	Dependent	660	9.65
Melpitz (DE0044K)	51.53	12.93	86	Dependent	670	8.78
Zugspitze-Schneefernerhaus (DE0054R)	47.42	10.98	2671	Independent	670	9.51
Leipzig-Eisenbahnstrasse (DE0066K)	51.35	12.41	120	Independent	670	9.51
Izaña (ES0018G)	28.41	-16.50	2373	Dependent	670	9.51
Granada (ES0020U)	37.16	-3.61	680	Dependent	670	9.51
Montsec (ES0022R)	42.05	0.73	1571	Dependent	670	9.51
El Arenosillo (ES0100R)	37.10	-6.73	41	Dependent	652	13.64
Montseny (ES1778R)	41.77	2.35	700	Dependent	670	8.48
Pallas (FI0096G)	67.97	24.12	565	Dependent	637	10.00
Observatoire Perenne de l' Environnement (FR0022R)	48.56	5.51	392	Dependent	880	7.24
Puy de Dôme (FR0030R)	45.77	2.96	1465	Dependent	670	9.51
Ispira (IT0004R)	45.80	8.63	209	Dependent	880	6.96
Mt Cimone (IT0009R)	44.18	10.70	2165	Dependent	670	9.51
Birkenes II (NO0002R)	58.39	8.25	219	Dependent	660	7.59
Zeppelin mountain (NO0042G)	78.91	11.89	474	Dependent	880	7.24

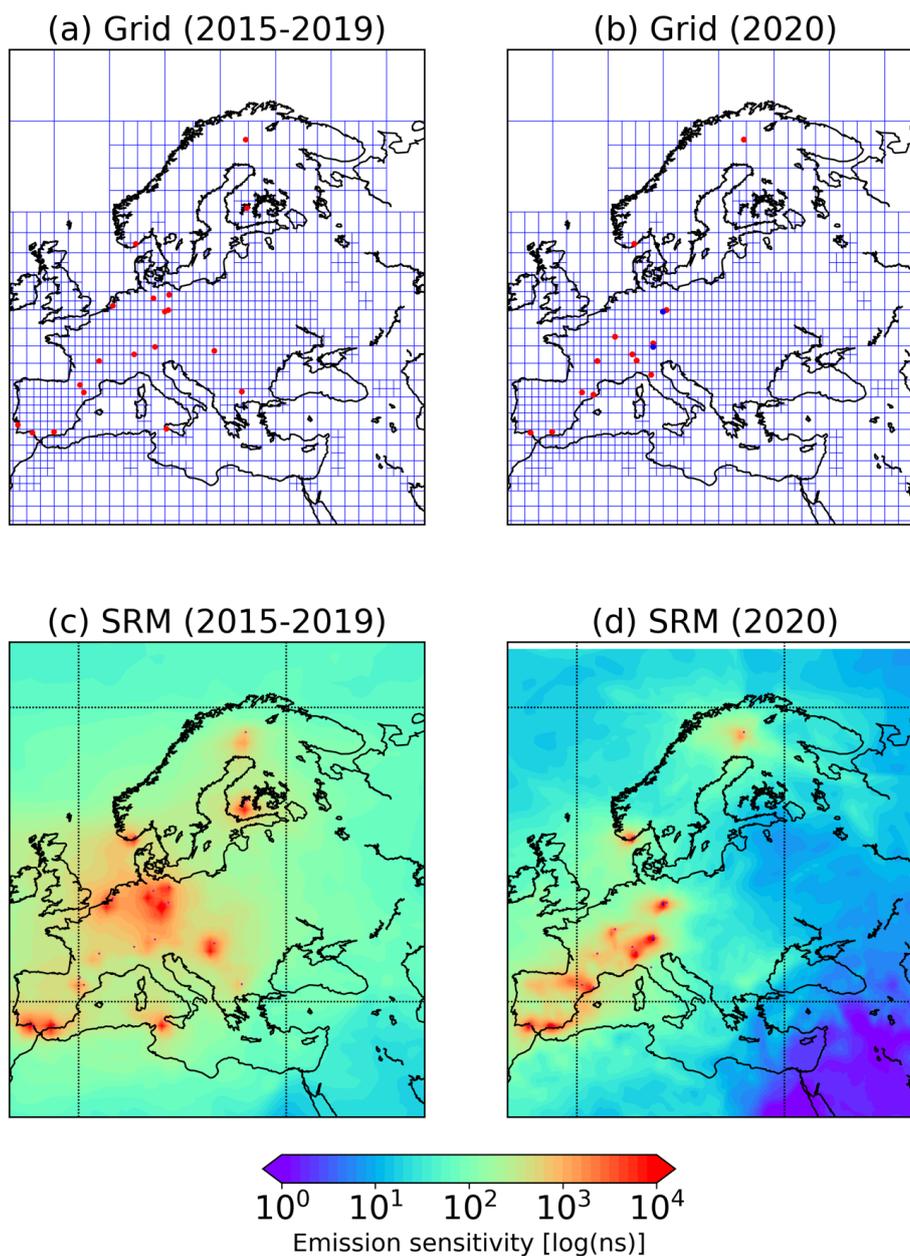
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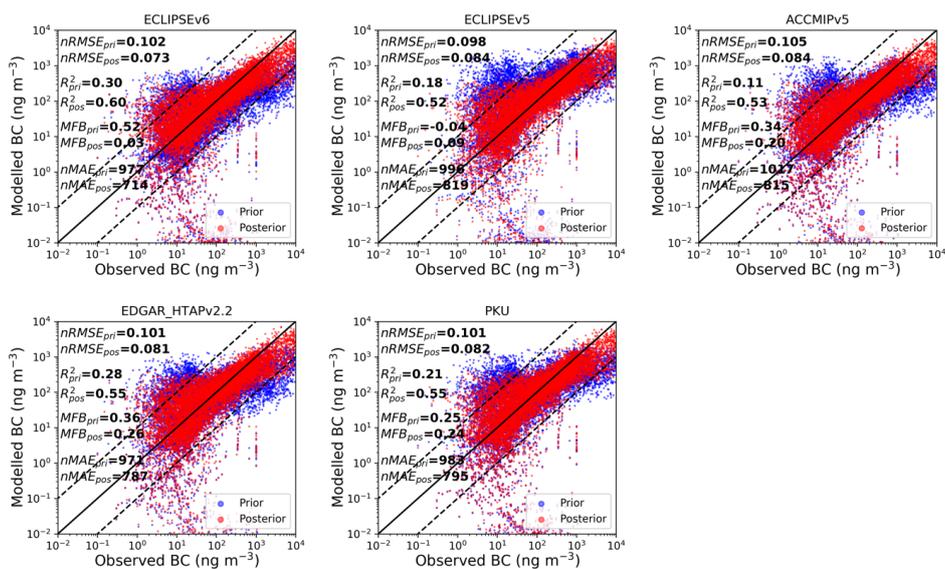
735 **Table 2.** Statistical measures (*RMSE*, Pearson's R^2 , *MFB* and *nMAE*) for each of the prior
736 and posterior concentrations against dependent observations (observations that were used in the
737 inversion algorithm) for BC (eBC). Note that the inversion using ECLIPSEv6 prior emission
738 dataset gave the best agreement with the observations and therefore the results of this inversion
739 are presented here.

	<i>nRMSE</i>	Pearson's R^2	<i>MFB</i>	<i>nMAE</i>
Prior ECLIPSEv6	0.102	0.30	0.52	997
Prior ECLIPSEv5	0.098	0.18	-0.04	996
Prior EDGAR_HTAPv2.2	0.105	0.11	0.34	1017
Prior ACCMIPv5	0.101	0.28	0.36	971
Prior PKU	0.101	0.21	0.25	983
Posterior ECLIPSEv6	0.073	0.60	0.03	714
Posterior ECLIPSEv5	0.084	0.52	0.09	819
Posterior EDGAR_HTAPv2.2	0.084	0.53	0.20	815
Posterior ACCMIPv5	0.091	0.55	0.26	787
Posterior PKU	0.082	0.55	0.24	795

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743 **Figure 1.** Aggregated inversion grid used for the (a) 2015–2019 and (b) 2020 inversions,
744 respectively. The dependent measurements that were used in the inversion were taken from
745 stations highlighted in red. The two independent stations that were used for the validation are
746 shown in blue. (c, d) Footprint emission sensitivity (i.e. SRM) averaged over all observations
747 and time steps for each of the inversions. Red points denote the location of each measurement
748 site.



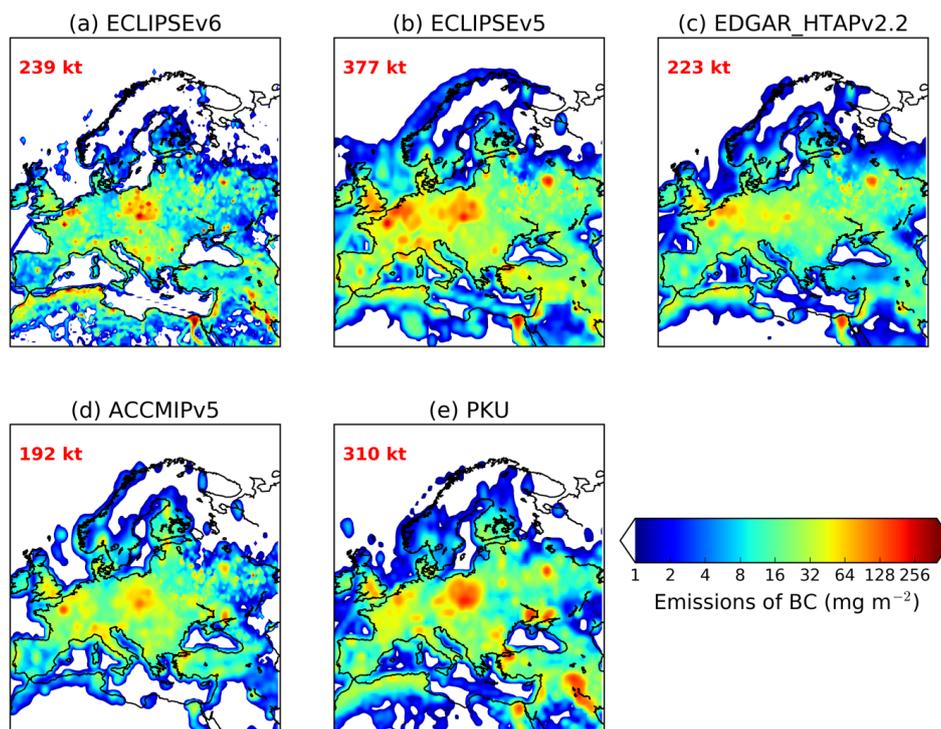
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750 **Figure 2.** Scatter plots of prior and posterior concentrations against dependent observations
751 (observations that were included in the inversion framework) from ACTRIS from January to
752 April 2020. Four statistical measures ($nRMSE$, Pearson's R^2 , MFB and $nMAE$) were used to
753 assess the performance of each inversion using five different prior emission inventories for BC
754 (ECLIPSEv5, v6, ACCMIPv5, EDGAR_HTAPv2.2 and PKU).

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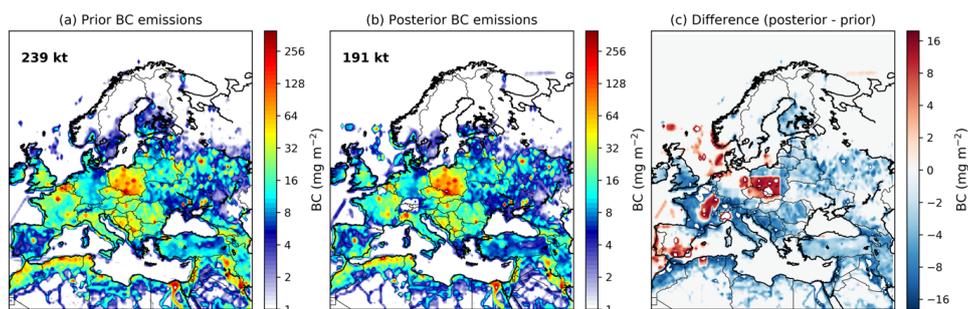


PRIOR EMISSIONS (JAN-APR 2020)



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757 **Figure 3.** Prior emissions of black carbon (BC) used in the inversions. BC emissions from
758 anthropogenic sources were adopted from ECLIPSE version 5 and 6 (Evaluating the CLimate
759 and Air Quality ImPacts of ShortlivEd Pollutants) (Klimont et al., 2017), EDGAR (Emissions
760 Database for Global Atmospheric Research) version HTAP_v2.2 (Janssens-Maenhout et al.,
761 2015), ACCMIP (Emissions for Atmospheric Chemistry and Climate Model Intercomparison
762 Project) version 5 (Lamarque et al., 2013) and PKU (Peking University) (Wang et al., 2014b).
763 Biomass burning emissions of BC from Global Fire Emissions Database (GFED) version 4.1
764 (Giglio et al., 2013) were added in each of the aforementioned inventories.

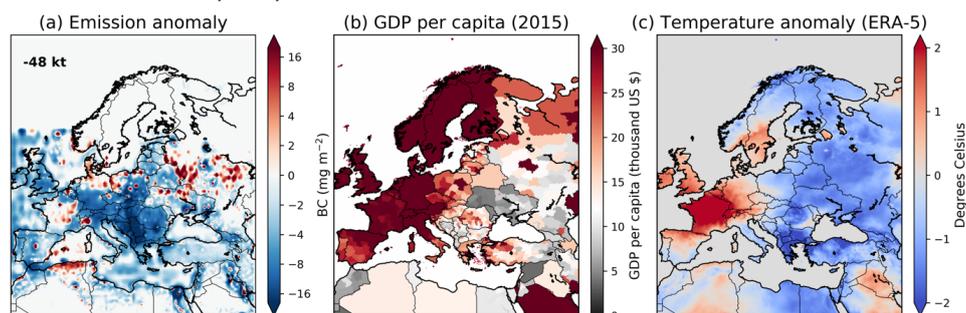
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767 **Figure 4.** (a) Prior emissions of BC from ECLIPSEv6, (b) optimized (posterior) BC emissions
768 after processing the ACTRIS data into the inversion algorithm, and (c) difference between
769 posterior and prior emissions. All the results correspond to the inversion yielding the best results
770 (**Table 2** and **Figure 2**).
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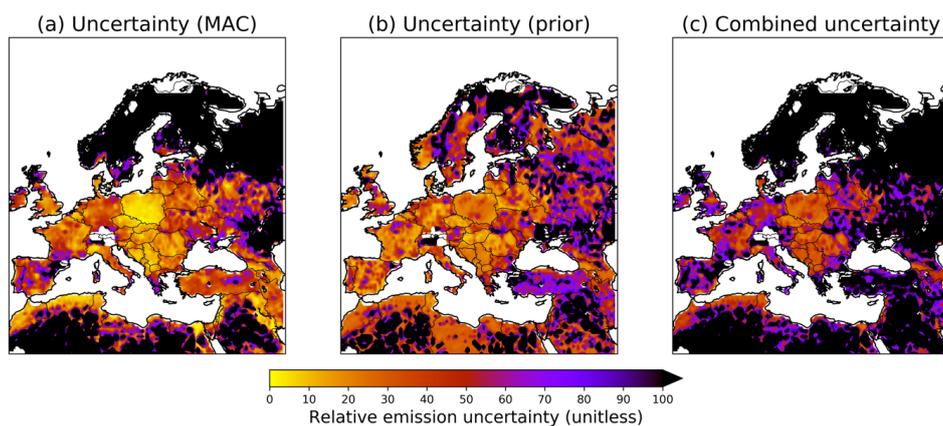


EMISSION ANOMALY, GDP, TEMPERATURE ANOMALY DURING THE LOCKDOWN PERIOD

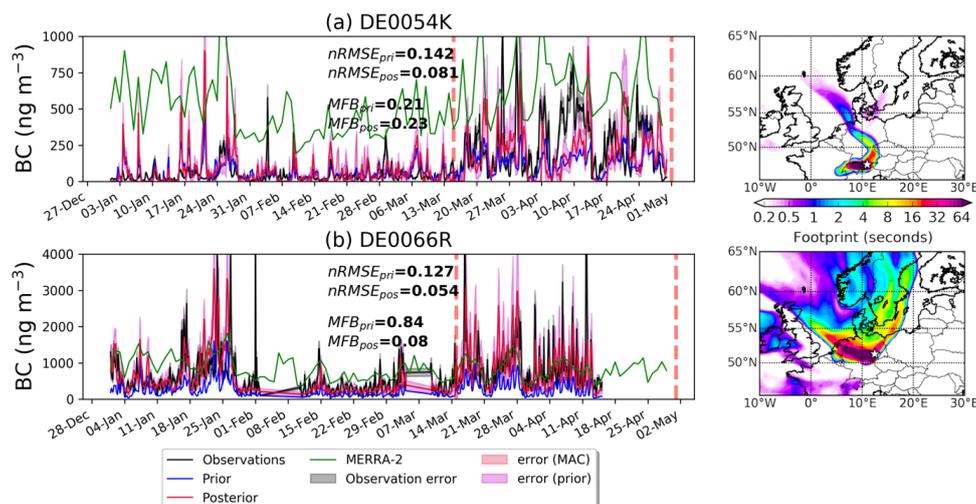


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773 **Figure 5.** (a) Difference in posterior BC emissions during the lockdown (14 March to 30 April
774 2020) in Europe from the respective emissions during the same period in 2015 – 2019, (b) GDP
775 from Kummu et al. (2020), and (c) temperature anomaly from ERA-5 (Copernicus Climate
776 Change Service (C3S), 2020) for the same period as the emission anomaly. The base GDP value
777 below which a low income can be assumed was set to 12 thousand US dollars.

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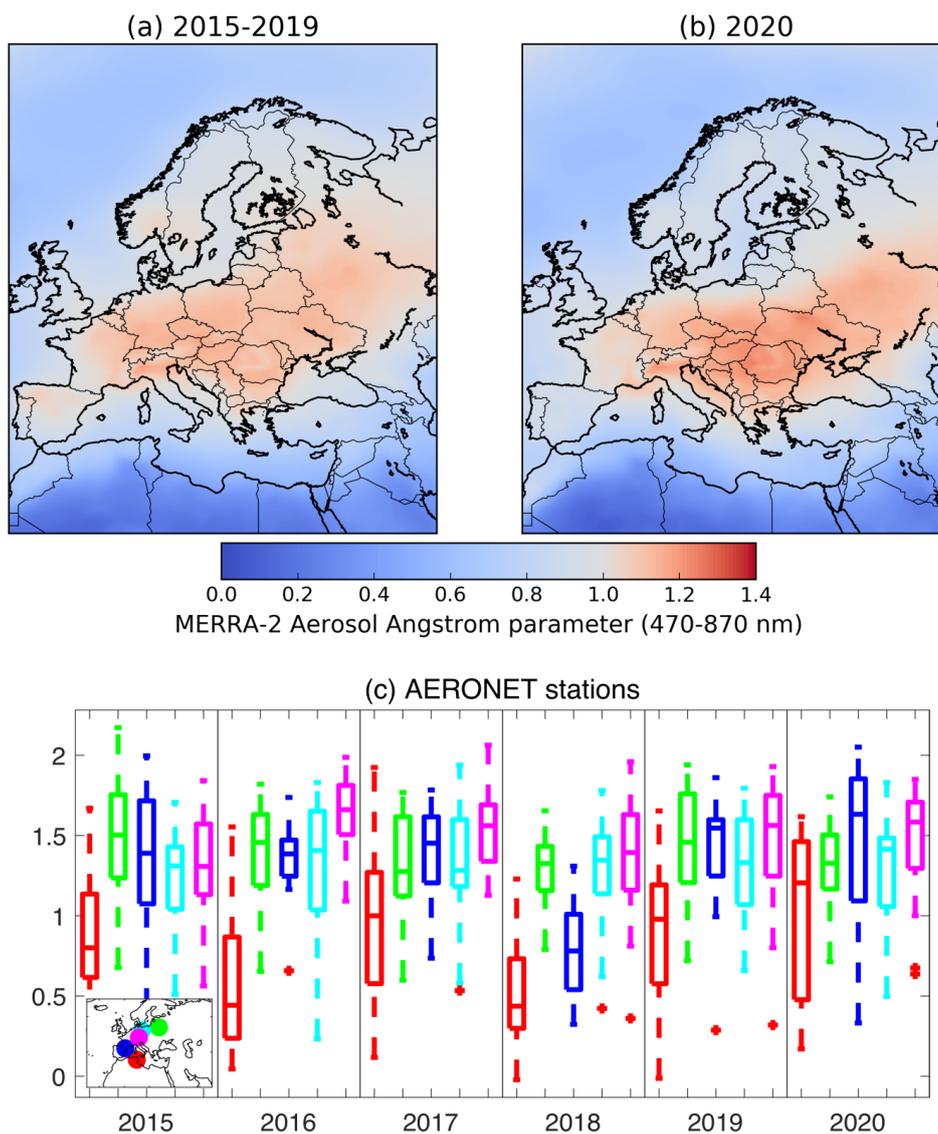


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780 **Figure 6.** (a) Uncertainty of BC emissions due to the use of variable MAC values to convert
781 from aerosol absorption to eBC concentrations that are used by the inversion algorithm. (b)
782 Uncertainty due to the use of five different prior emissions inventories for BC. (c) Combined
783 uncertainty.
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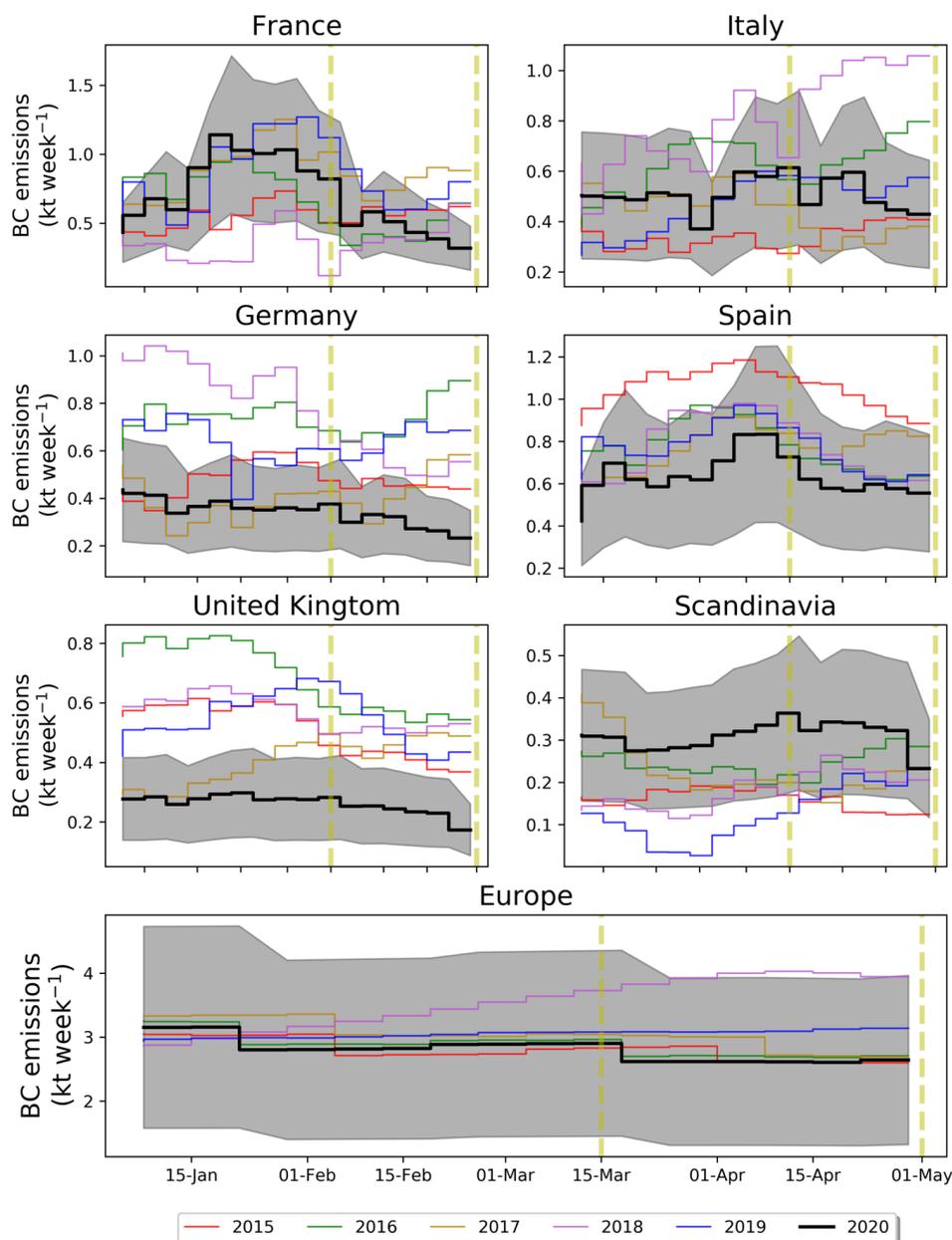
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Figure 7. Prior and posterior BC concentrations at DE0054K and DE0066R stations that were not included in the inversion are compared with observations. The validation is done by calculating the *nRMSEs* and *MFBS* for the prior and posterior concentrations. The uncertainty of the observations is also given together with the posterior uncertainties in the concentrations calculated from the use of different MAC and prior emissions. For comparison, we plot the concentrations from MERRA-2 at the same two stations. The vertical dashed lines denote the period of the lockdown in most of Europe. On the right, the average footprint emission sensitivities are given at each independent station for the period of the lockdown.



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Figure 8. (a) Average total aerosol Ångström parameter (470-870 nm) over Europe (mid-March to April) in the five previous years (2015–2019) and (b) in 2020 (lockdown). (c) AERONET Absorption AE in Ben Salem (9.91°E, 35.55°N), Minsk (27.60°E, 53.92°N), Montsec (0.73°E, 42.05°N), MetObs Lindenberg (14.12°E, 52.21°N) and Munich University (11.57°E, 48.15°N) during mid-March to April in all years since 2015.



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803 **Figure 9.** Posterior BC emissions in the most highly affected European countries (France, Italy,
804 Germany, Spain and UK), Scandinavia and Europe from the COVID-19 pandemic (2020).
805 Posterior BC emissions for every year since 2015 are also plotted in the same temporal
806 resolution to show changes in BC emissions characteristics during the 2020 COVID-19
807 pandemic. The grey shaded area corresponds to the BC emission uncertainty, while the vertical
808 yellow dashed lines correspond to the beginning and end of the 2020 lockdown.