



22 from open websites were applied to test the transferring ability of the novel model, and confirmed



that the model was robust to predict the surface 8-h O<sub>3</sub> concentration during other periods ( $R^2 = 0.67$ , RMSE = 25.68  $\mu\text{g}/\text{m}^3$ ). RF-GAM was then used to predict the daily 8-h O<sub>3</sub> level over Tibetan Plateau during 2005-2018 for the first time. It was found that the estimated O<sub>3</sub> concentration displayed a slow increase from  $64.74 \pm 8.30 \mu\text{g}/\text{m}^3$  to  $66.45 \pm 8.67 \mu\text{g}/\text{m}^3$  from 2005 to 2015, whereas it decreased from the peak to  $65.87 \pm 8.52 \mu\text{g}/\text{m}^3$  during 2015-2018. Besides, the estimated 8-h O<sub>3</sub> concentrations exhibited notably spatial variation with the highest values in some cities of North Tibetan Plateau such as Huangnan ( $73.48 \pm 4.53 \mu\text{g}/\text{m}^3$ ) and Hainan ( $72.24 \pm 5.34 \mu\text{g}/\text{m}^3$ ), followed by the cities in the central region including Lhasa ( $65.99 \pm 7.24 \mu\text{g}/\text{m}^3$ ) and Shigatse ( $65.15 \pm 6.14 \mu\text{g}/\text{m}^3$ ), and the lowest one in a city of Southeast Tibetan Plateau named Aba ( $55.17 \pm 12.77 \mu\text{g}/\text{m}^3$ ). Based on the 8-h O<sub>3</sub> critical value (100  $\mu\text{g}/\text{m}^3$ ) scheduled by World Health Organization (WHO), we further estimated the annually mean nonattainment days over Tibetan Plateau. It should be noted that most of the cities in Tibetan Plateau shared with the excellent air quality, while several cities (e.g., Huangnan, Haidong, and Guoluo) still suffered from more than 40 nonattainment days each year, which should be paid more attention to alleviate local O<sub>3</sub> pollution. The result shown herein confirms the novel hybrid model improves the prediction accuracy and can be applied to assess the potential health risk, particularly in the remote regions with sparse monitoring sites.

**Keywords:** Surface O<sub>3</sub> level; satellite data; random forest; generalized additive model; Tibetan Plateau

## 1. Introduction

Along with the rapid economic development and urbanization, the anthropogenic emissions of nitrogen oxides (NO<sub>x</sub>) and volatile organic compounds (VOCs) displayed high-speed growth. The chemical reaction between NO<sub>x</sub> and VOCs in the presence of sunlight was beneficial to the ambient



45 ozone ( $\text{O}_3$ ) formation (Wang et al., 2019; Wang et al., 2017). As a strong oxidant, ambient  $\text{O}_3$  could  
46 play a negative role on human health through aggravating the cardiovascular and respiratory  
47 function (Ghude et al., 2016; Marco, 2017; Yin et al., 2017a). Apart from the effect on human health,  
48  $\text{O}_3$  also posed a great threaten on vegetation growth (Emberson, 2017; Feng et al., 2015; Qian et al.,  
49 2018; Feng et al., 2019). Moreover, the tropospheric  $\text{O}_3$  can perturb the radiative energy budget of  
50 the earth-atmosphere system as the third most important greenhouse gas next to carbon dioxide  
51 ( $\text{CO}_2$ ) and methane ( $\text{CH}_4$ ), thereby changing the global climate (Bornman et al., 2019; Fu et al.,  
52 2019; Wang et al., 2019). Recently, the particulate matter less than  $2.5\ \mu\text{m}$  ( $\text{PM}_{2.5}$ ) concentration  
53 showed the persistent decrease, while the  $\text{O}_3$  issue has been increasingly prominent in China (Li et  
54 al., 2017b; Li et al., 2019b). Therefore, it was critical to accurately reveal the spatiotemporal  
55 variation of  $\text{O}_3$  pollution and assess its health risk in China.

56 A growing body of studies began to investigate the spatiotemporal variation of  $\text{O}_3$  level  
57 worldwide. Wang et al. (2014b) demonstrated that the 8-h  $\text{O}_3$  concentration in nearly all of the  
58 provincial cities experienced the remarkable increase during 2013-2014. Following the work, Li et  
59 al. (2017) reported that the annually mean  $\text{O}_3$  concentration over China increased by 9.18% during  
60 2014-2016. In other Asian countries except China, Vellingiri et al. (2015) performed long-term  
61 observation and found that the  $\text{O}_3$  concentration in Seoul, South Korea displayed gradual increase in  
62 the past decades. In the Southeast United States, Li et al. (2018) observed that the surface  $\text{O}_3$   
63 concentration also displayed the gradual decrease in the recent ten years. Although the number of  
64 ground-level monitoring sites have been increasing globally, the limited monitoring sites still cannot  
65 accurately reflect the fine-scale  $\text{O}_3$  pollution status because each site shows small spatial  
66 representativeness ( $0.25\text{-}16.25\ \text{km}^2$ ) (Shi et al., 2018). Furthermore, the number of monitoring



67 sites in many countries (e.g., China and the United States) displays uneven distribution  
68 characteristic at the spatial scale. In China, most of these sites focus on North China Plain (NCP)  
69 and Yangtze River Delta (YRD), while West China extremely lacks of the ground-level O<sub>3</sub> data,  
70 which often increases the uncertainty of health assessment. Therefore, many studies used  
71 various models to estimate the O<sub>3</sub> concentration without monitoring sites. Chemical transport  
72 models (CTMs) were often considered as the typical methods to predict the surface O<sub>3</sub> level.  
73 Zhang et al. (2011) employed the Geos-Chem model to simulate the surface O<sub>3</sub> concentration  
74 over the United States, suggesting that the model could capture the spatiotemporal variation of  
75 surface O<sub>3</sub> concentration at a large spatial scale. Later on, Wang et al. (2016) developed a hybrid  
76 model named land use regression (LUR) coupled with CTMs to predict the surface O<sub>3</sub>  
77 concentration in the Los Angeles Basin, California. In recent years, these methods were also  
78 applied to estimate the surface O<sub>3</sub> level over China. Liu et al. (2018) used Community  
79 Multiscale Air Quality (CMAQ) model to simulate the nationwide O<sub>3</sub> concentration over China  
80 in 2015. Nonetheless, the high-resolution O<sub>3</sub> prediction using CTMs might be widely deviated  
81 from the measured value owing to the imperfect knowledge about the chemical mechanism and  
82 the higher uncertainty of emission inventory. Moreover, the continuous emission data of NO<sub>x</sub>  
83 and VOCs were not always open access, which restricted the long-term estimation of surface  
84 O<sub>3</sub> concentration using CTMs.

85 Fortunately, the daily satellite data enables the fine-scale estimation of O<sub>3</sub> level at a regional  
86 scale due to broad spatial coverage and high temporal resolution (McPeters et al., 2015). Shen  
87 et al. (2019) confirmed that satellite retrieved O<sub>3</sub> column amount can accurately reflect the  
88 spatiotemporal distribution of surface O<sub>3</sub> level. Therefore, some studies tried to use traditional



89 statistical models coupled with high-resolution satellite data to estimate the ambient O<sub>3</sub> level.  
90 Fioletov et al. (2002) used the satellite measurement to investigate the global distribution of O<sub>3</sub>  
91 concentration based on simple linear model. Recently, Kim et al. (2018) employed the  
92 integrated empirical geographic regression method to predict the long-term (1979-2015)  
93 variation of ambient O<sub>3</sub> concentration over United States based on O<sub>3</sub> column amount data.  
94 Although the statistical modelling of ambient O<sub>3</sub> concentration is widespread all around the  
95 world, most of these traditional statistical modelling only utilized the linear model to predict  
96 the ambient O<sub>3</sub> concentration, which generally decreased the prediction performance because  
97 the nonlinearity and high-order interactions between O<sub>3</sub> and predictors cannot be managed by  
98 a simple linear model.

99 As an extension of traditional statistical model, machine learning methods have been widely  
100 applied to estimate the pollutant level in recent years because of their excellent predictive  
101 performances. Among these machine learning algorithms, decision tree models such as random  
102 forest (RF) and extreme gradient boosting (XGBoost) strike a perfect balance between  
103 prediction performance and computing cost. Furthermore, decision tree models can obtain the  
104 contribution of each predictor to air pollutant, which was beneficial to the parameter adaption  
105 and model optimization. Chen et al. (2018b) firstly employed RF model to simulate the PM<sub>2.5</sub>  
106 level in China since 2005. Following this work, we recently used the XGBoost model to  
107 estimate the 8-h O<sub>3</sub> concentration in Hainan Island for the first time and captured the moderate  
108 predictive performance ( $R^2 = 0.59$ ) (Li et al., 2020). While decision tree model shows many  
109 advantages in predicting pollutant level, the spatiotemporal autocorrelation of pollutant  
110 concentration is not concerned by these studies. Li et al. (2019a) confirmed that the prediction



error by decision tree model varied greatly with space and time. Thus, it is imperative to incorporate the spatiotemporal variables into the original model to further improve the performance. To resolve the defects of decision tree models, Zhan et al. (2018) developed a hybrid model named RF-spatiotemporal Kriging (STK) to predict the  $O_3$  concentration over China and achieved the better performance (Overall  $R^2 = 0.69$ , Southwest China  $R^2 = 0.66$ ). Unfortunately, RF-STK model still showed some weaknesses in predicting  $O_3$  concentration. First of all, the predictive performance of the STK model was strongly dependent on the number of monitoring sites and their spatial density. The model often showed worse predictive performance in the region with sparse monitoring sites (Gao et al., 2016). Moreover, the ensemble model cannot simulate the  $O_3$  level during the periods without ground-measured data. In contrast, generalized additive model (GAM) not only considers the time autocorrelation of  $O_3$  concentration, but also shows the better extrapolation ability (Chen et al., 2018a; Ma et al., 2015). Thus, the ensemble model of RF and GAM is proposed to predict the spatiotemporal variation of surface 8-h  $O_3$  concentration.

Tibetan Plateau, the highest plateau around the world, shows the higher surface solar radiation compared with the region outside the plateau. It was well documented that high solar radiation is beneficial to generate large amount of OH radical, resulting in the  $O_3$  formation via the reaction of VOC and OH radical (Ou et al., 2015). While the total  $O_3$  column amount in Tibetan Plateau displayed the slight decrease since 1990s, the convergent airflow formed by subtropical anticyclones could bring ozone-rich air surrounding the plateau to the low atmosphere (Lin et al., 2008), thereby leading to the higher surface  $O_3$  concentration over the plateau. Most studies focused on the stratosphere-troposphere transport of  $O_3$  in Tibetan Plateau,



133 whereas limited effort was spared to investigate ground-level O<sub>3</sub> level over this region. To date,  
134 only several studies concerned about the spatiotemporal variation of surface O<sub>3</sub> concentration  
135 in this region (Chen et al., 2019; Shen et al., 2014; Yin et al., 2017b). Furthermore, some of  
136 these field-observation studies only used the limited monitoring sites to reveal the  
137 spatiotemporal variation of O<sub>3</sub> concentration, while they cannot clarify the real O<sub>3</sub> status in  
138 many regions without monitoring sites (e.g., Northern part of Tibetan Plateau). Apart from these  
139 field measurements, Liu et al. (2018) ( $R = 0.60$ ) and Zhan et al. (2018) ( $R^2 = 0.66$ ) used CTM  
140 and machine learning model to simulate the surface O<sub>3</sub> concentration over China in 2015,  
141 respectively. Both of these studies included the predicted O<sub>3</sub> level in Tibetan Plateau. Although  
142 they have finished the pioneering work, the predictive performances of both studies were not  
143 very excellent. Therefore, it was imperative to develop a higher quality model to enhance the  
144 modelling accuracy.

145 Here, we developed a new hybrid method (RF-GAM) model integrating satellite data,  
146 meteorological factors, and geographical variables to simulate the gridded 8-h O<sub>3</sub> concentration  
147 over Tibetan Plateau for the first time. Based on the estimated surface O<sub>3</sub> concentration, we  
148 clarified the long-term variation (2005-2018) of surface O<sub>3</sub> concentration and quantified the  
149 key factors for the annual trend. Filling the gap of statistical estimation 8-h O<sub>3</sub> level in a remote  
150 region, this study provides useful datasets for epidemiological studies and air quality  
151 management.

## 152 2. Materials and methods

### 153 2.1 Study area

154 Tibetan Plateau is located in Southwest China ranging from 26.00 to 39.58°N and from



155 73.33 to 104.78°E. Tibetan Plateau is surrounded by Taklamakan Desert to north, Sichuan Basin  
156 to southeast. The land area of Tibetan Plateau reaches 2.50 million km<sup>2</sup> (Chan et al., 2006).  
157 Based on the air circulation pattern, Tibetan Plateau can be roughly classified into the monsoon-  
158 influenced region and the westerly-wind influenced region (Wang et al., 2014a). The annually  
159 mean air temperature in most regions are below 0°C. The annually mean rainfall amount in  
160 Tibetan Plateau ranges from 50 to 2000 mm. The terrain conditions are complex and the higher  
161 altitude focused on the central region. Tibetan Plateau is generally treated as the remote region  
162 lack of anthropogenic activity and most of the residents focus on southeast and south part of  
163 Tibetan Plateau. Tibetan Plateau is consisted of 19 prefecture-level cities and their names and  
164 corresponding geographical locations are shown in Fig. 1 and Fig. S1.

## 165 2.2 Data preparation

### 166 2.2.1 Ground-level 8-h O<sub>3</sub> concentration

167 The daily 8-h O<sub>3</sub> data in 37 monitoring sites over Tibetan Plateau from May 13th, 2014 to  
168 December 31th, 2018 were collected from the national air quality monitoring network. The O<sub>3</sub> levels  
169 in all of these sites were determined using an ultraviolet-spectrophotometry method. The highest 8-  
170 h moving average O<sub>3</sub> concentration each day was calculated as the daily 8-h O<sub>3</sub> level after data  
171 quality assurance. The data quality of all the monitoring sites were assured on the basis of the HJ  
172 630-2011 specifications. The data with no more than two consecutive hourly measurement missing  
173 in each day was treated as the valid data.

### 174 2.2.2 Satellite-retrieved O<sub>3</sub> column amount

175 The O<sub>3</sub> column amount (DU) during 2005-2018 were downloaded from the Ozone Monitoring  
176 Instrument-O<sub>3</sub> (OMI-O<sub>3</sub>) level-3 data with a 0.25° spatial resolution from the website of National



177 Aeronautics and Space Administration (NASA) (<https://www.nasa.gov/>). The OMI-O<sub>3</sub> product  
178 shows global coverage and traverses the earth once a day. The O<sub>3</sub> column amount with cloud  
179 radiance fraction > 0.5, terrain reflectivity > 30%, and solar zenith angles > 85° should be removed.  
180 In addition, the cross-track pixels significantly influenced by row anomaly should be deleted.

### 181 2.2.3 Meteorological data and geographical covariates

182 The daily meteorological data were obtained from ERA-Interim datasets with 0.125° resolution.  
183 These meteorological data were consisted of 2 meter dewpoint temperature (d2m), 2 meter  
184 temperature (t2m), 10 meter U wind component (u10), 10 meter V wind component (v10), boundary  
185 layer height (blh), sunshine duration (sund), surface pressure (sp), and total precipitation (tp). The  
186 30 m-resolution elevation data (DEM) was downloaded from China Resource and Environmental  
187 Science Data Center (CRESDC). The data of gross domestic production (GDP) and population  
188 density with 1 km resolution were also extracted from CRESDC. Population density and GDP in  
189 2005, 2010, and 2015 were integrated into the model to predict the surface 8-h O<sub>3</sub> concentration  
190 over Tibetan Plateau because these data were available each five years. Additionally, the land use  
191 data of 30 m resolution (e.g., waters, grassland, urban, forest) were also extracted from CRESDC.  
192 At last, the latitude, longitude, and time were also incorporated into the model.

193 All of the explanatory variables collected were resampled to 0.25° × 0.25° grids to predict the  
194 O<sub>3</sub> level. The original meteorological data with 0.125° resolution were resampled to 0.25° grid. The  
195 land use area, elevation, GDP and population density in each grid were calculated using spatial  
196 clipping. Lastly, all of the predictors were integrated into an intact table to train the model.

### 197 2.3 Model development and assessment

198 The RF-GAM model was regarded as the hybrid model of RF and GAM. The RF-GAM model



is a two-stage model that the prediction error estimated by the RF model was then simulated by GAM. The prediction results of RF and GAM were summed as the final result of RF-GAM model (Fig. 2). The detailed equation is as follows:

$$Z(s, t) = P(s, t) + E(s, t) \quad (1)$$

where  $Z(s, t)$  is the estimated 8-h  $O_3$  level at the location  $s$  and time  $t$ ;  $P(s, t)$  represents the 8-h  $O_3$  concentration predicted by the RF model;  $E(s, t)$  denotes the prediction error by GAM.

In the RF model, a large number of decision trees were planted based on the bootstrap sampling method. At each node of the decision tree, a random sample of all predictors was applied to determine the best split among them. Following the procedure, a simple majority vote was employed to predict the 8-h  $O_3$  level. The RF model avoided priori linear assumption of  $O_3$  concentration and predictors, which was often not in good agreement with actual state. The RF model has two key parameters including  $n_{tree}$  (the number of trees grown) and  $m_{try}$  (the number of explanatory variables sampled for splitting at each node). The prediction performance of the RF model was strongly dependent on the two parameters. The optimal  $n_{tree}$  and  $m_{try}$  were determined based on the least out-of-bag (OOB) errors. Besides, the backward variable selection method was performed on the RF submodel to achieve the better performance. At each step of the predictor selection, the variable with the least important value was excluded from the next step. This one-variable-at-a-time exclusion method was repeated until only two explanatory variables remained in the submodel. Finally, all of the selected variables except the area of waters were integrated into the model to achieve the best prediction performance. The detailed RF model is as follows:

$$O_3 = O_3 \text{ column} + \text{Elevation} + \text{Agr} + \text{Urban} + \text{Forest} + \text{GDP} + \text{Grassland} + \text{Population} + \text{Prec} + T + WS + P + \text{tsun} + RH \quad (2)$$



220 where  $O_3$  denotes the observed 8-h  $O_3$  level in the monitoring site; the  $O_3$  column represents the  $O_3$   
221 column amount in the corresponding grid; Elevation denotes the corresponding elevation of the site;  
222 Agr, Urban, Forest, Grassland are the agricultural land, urban land, forest land, and the grassland,  
223 respectively. Population represents the population density in the corresponding site. Prec, T, WS, P,  
224 tsun, and RH are precipitation, air temperature, wind speed, air pressure, sunshine duration, and  
225 relative humidity, respectively. Additionally, other five models including RF, generalized regression  
226 neutral network (GRNN), backward propagation neural network (BPNN), Elman neural network  
227 (ElmanNN), and extreme learning machine (ELM) also used the backward variable selection  
228 method. The  $R^2$  value was treated as an important parameter to add or reduce the variable. The  
229 variable should be removed when the  $R^2$  value of the submodel showed the remarkable decrease  
230 with the integration of this variable. Lastly, the optimal variable group was applied to establish the  
231 submodel.

232 Following the RF submodel, the prediction error estimated by the RF submodel was further  
233 modelled by the GAM. GAM could reflect the time autocorrelation of predictive error of RF model,  
234 and thus the ensemble model of RF and GAM might decrease the modelling error of one-stage  
235 model. All of the variables were incorporated into the models to establish the second-stage model,  
236 and the backward variable selection was also used to determine the optimal variable group.

237 The 10-fold cross-validation (CV) technique was employed to evaluate the predictive  
238 performances for all of the machine learning models. All of the training data set were randomly  
239 classified into 10 subsets uniformly. In each round of validation, nine subsets were used to train and  
240 the remaining subset was applied to test the model performance. The process was repeated 10 times  
241 until every subset has been tested. Some statistical indicators including  $R^2$ , Root Mean Square Error



(RMSE), Mean Prediction Error (MPE) and the slope were calculated to assess the model performance. The optimal model with the best performance was used to estimate the 8-h O<sub>3</sub> concentration in the past decades.

### 3. Results and discussion

#### 3.1 The validation of model performance

Figure 3 shows the density scatterplots of the fitting and 10-fold cross-validation results for eight machine learning models for China. The 10-fold cross-validation R<sup>2</sup> values followed the order of RF-GAM (R<sup>2</sup> = 0.76) > RF-STK (R<sup>2</sup> = 0.63) > RF (R<sup>2</sup> = 0.55) > GRNN (R<sup>2</sup> = 0.53) > BPNN (R<sup>2</sup> = 0.50) > XGBoost (R<sup>2</sup> = 0.48) > ElmanNN (R<sup>2</sup> = 0.47) > ELM (R<sup>2</sup> = 0.32). The RMSE values of RF-GAM, RF-STK, RF, GRNN, XGBoost, BPNN, ElmanNN, and ELM were 14.41, 17.79, 19.13, 19.41, 20.73, 20.06, 20.61, and 23.36 µg/m<sup>3</sup>, respectively. MPE showed the similar characteristic with RMSE in the order of RF-GAM (10.97 µg/m<sup>3</sup>) < RF-STK (13.48 µg/m<sup>3</sup>) < RF (14.71 µg/m<sup>3</sup>) < GRNN (14.89 µg/m<sup>3</sup>) < BPNN (15.43 µg/m<sup>3</sup>) < ElmanNN (15.75 µg/m<sup>3</sup>) < XGBoost (15.80 µg/m<sup>3</sup>) < ELM (18.23 µg/m<sup>3</sup>). Besides, the slope of the RF-GAM model was closer to 1 compared with other models. It was well documented that the RF model generally showed the better performance than other models because this method did not need to define complex relationships between the explanatory variables and the O<sub>3</sub> concentration (e.g., linear or nonlinear). Furthermore, the variable importance indicators calculated by the RF model can help user to distinguish the key variables from noise ones and make full use of the strength of each predictor to assure the model robustness. Although BPNN, GRNN, XGBoost, ElmanNN, and ELM have been widely applied to estimate the air pollutant concentrations (Chen et al., 2018c; Zang et al., 2018; Zhu et al., 2019), these methods suffered from some weaknesses in predicting the pollutant level. For instance, both of BPNN and



ElmanNN models could capture the locally optimal solution when the training subsets were integrated into the final model, which decreased the predictive performance of the model (Wang et al., 2015). Moreover, BPNN generally showed slow training speed, especially with the huge training subsets (Li and Park, 2009; Wang et al., 2015). ELM often consumed more computing resource and experienced the over-fitting issue due to the increase of sampling size (Huang et al., 2015; Shao et al., 2015). GRNN method advanced the training speed compared with BPNN model and avoided the locally optimal solution during the modelling process (Zang et al., 2019), whereas the predictive performance is still worse than that of RF model. XGBoost was often considered to be robust in predicting air pollutant level (Li et al., 2020), while the model did not display the excellent performance in the present study. It might be attributable to that the sampling size in the present study was not enough because the model generally showed the better performance with big samples. Moreover, we found that the two-stage model was superior to the one-way model in the predictive performance. The encouraging result suggested that the relationship between the predictors and the 8-h  $O_3$  concentration varied with space and time. The two-stage model used the GAM method to further adjust the prediction error of the RF model, and considered the spatiotemporal correlation of predictor error in Tibetan Plateau. Although the STK model incorporated space and time into the model simultaneously, the RF-GAM model outperformed the RF-STK model. It was assumed that the STK model showed the higher uncertainty in predicting the  $O_3$  concentration in the region with scarce sampling sites (Gao et al., 2016; Li et al., 2017a). Overall, the ensemble RF-GAM model showed the significant improvement in predictive performance.

The performances of the RF-GAM model in four seasons were also assessed by 10-fold cross-validation (Tab. 1). The predictive performance of the RF-GAM model showed significantly



seasonal difference with the highest  $R^2$  value observed in summer (0.74), followed by winter (0.69) and autumn (0.67), and the lowest one in spring (0.64). However, both of RMSE and MPE displayed different seasonal characteristics with the  $R^2$  value. Both of RMSE and MPE for RF-GAM followed the order of spring ( $15.32$  and  $11.94 \mu\text{g}/\text{m}^3$ ) > summer ( $15.13$  and  $11.75 \mu\text{g}/\text{m}^3$ ) > winter ( $14.58$  and  $11.44 \mu\text{g}/\text{m}^3$ ) > autumn ( $13.23$  and  $10.52 \mu\text{g}/\text{m}^3$ ). The lowest  $R^2$  value in spring might be caused by multiple  $\text{O}_3$  sources and complicate  $\text{O}_3$  formation mechanisms. The large estimation errors (e.g., RMSE and MPE) in spring and summer were attributable to the high 8-h  $\text{O}_3$  concentration in these seasons, while the low prediction error observed in autumn was contributed by the low  $\text{O}_3$  level.

Apart from the seasonal variation, we also investigated the spatial variability of the predictive accuracy for RF-GAM model. Tibetan Plateau was classified into five provinces and then the predictive performance of RF-GAM model in each province was calculated. Among the five provinces, Gansu province displayed the highest  $R^2$  value (0.74), followed by Sichuan province (0.71), Qinghai province (0.70), Tibet autonomous region (0.69), and Yunnan province (0.54) (Tab. 2). The result shown herein was not in agreement with the previous studies by Geng et al. (2018), who confirmed that the predictive performance of machine learning model was positively associated with the sampling size. It was assumed that the spatial distribution of the sampling sites in Tibet was uneven and the sampling density is low, though Tibet possessed the maximum monitoring sites compared with other provinces. The prediction error (RMSE and MPE) did not exhibit the same characteristics with the  $R^2$  value. The higher RMSE and MPE focused on Tibet autonomous region ( $14.81$  and  $11.24 \mu\text{g}/\text{m}^3$ ) and Qinghai province ( $14.83$  and  $11.33 \mu\text{g}/\text{m}^3$ ) due to the higher values of blh and sund. The lowest values of RMSE and MPE could be observed in Yunnan province, which was contributed by the higher rainfall amount.



Although 10-fold cross-validation verified that the RF-GAM model showed the better predictive performance in estimating the surface 8-h  $O_3$  concentration, the test method cannot validate the transferring ability of the final model. The monitoring site in Tibetan Plateau before May, 2014 is very limited, and only the daily 8-h  $O_3$  data in Lhasa from the open website (<https://www.aqistudy.cn/historydata/>) was available to compare with the simulated data. As depicted in Fig. 4, the  $R^2$  value of unlearning 8-h  $O_3$  level against predicted 8-h  $O_3$  concentration reached 0.67, which was slightly lower than that of the 10-fold cross-validation  $R^2$  value. Overall, the extrapolation ability of the RF-GAM model is satisfactory, and thus it was supposed that the model could be applied to estimate the  $O_3$  concentration in other years. Both of RMSE and MPE for the unlearning 8-h  $O_3$  level against the predicted 8-h  $O_3$  concentration were significantly higher than those of the 10-fold cross-validation. It was supposed that Lhasa showed the higher surface 8-h  $O_3$  concentration over Tibetan Plateau.

### 3.2 Variable importance

The results of variable importance for key variables are depicted in Fig. 5. In the final RF-GAM model, it was found that time was the dominant factor for the 8-h  $O_3$  concentration in Tibetan Plateau, indicating that the ambient  $O_3$  concentration displayed significantly temporal correlation. Following the time, meteorological factors served as the main factors for the  $O_3$  pollution in the remote region. The sum of sund, sp, d2m, t2m, and tp occupied 34.43% of the overall variable importance. Among others, sund was considered to be the most important meteorological factors for the  $O_3$  pollution. It was assumed that strong solar radiation and long duration of sunshine favored the photochemical generation of ambient  $O_3$  (Malik and Tauler, 2015; Stähle et al., 2018). Tan et al. (2018) demonstrated that the chemical reaction between  $NO_x$  and VOCs was strongly dependent on the



330 sunlight. Besides, the atmospheric pressure (sp) was also treated as a major driver for the O<sub>3</sub>  
331 pollution over Tibetan Plateau. Santurtún et al. (2015) have demonstrated that sp was closely linked  
332 to the atmospheric circulation and synoptic scale meteorological pattern, which could influence the  
333 long-range transport of ambient O<sub>3</sub>. Apart from sund and sp, d2m and t2m played significant role  
334 on the O<sub>3</sub> pollution, which was in consistent with many previous studies (Zhan et al., 2018). Zhan  
335 et al. (2018) observed that cold temperature was not favorable to the O<sub>3</sub> formation. d2m can affect  
336 the surface O<sub>3</sub> pollution through two aspects. On the one hand, RH affected heterogeneous reactions  
337 of O<sub>3</sub> and particles (e.g., soot, mineral) (He et al., 2017; He and Zhang, 2019). On the other hand,  
338 high RH could increase the soil moisture and evaporation, and thus the water-stressed plants tended  
339 to emit more biogenic isoprene, thereby promoting the elevation of O<sub>3</sub> concentration (Zhang and  
340 Wang, 2016). It should be noted that the effect of precipitation on O<sub>3</sub> pollution was relatively weaker  
341 than those of other meteorological factors. Zhan et al. (2018) also found the similar result and  
342 believed that rain scavenging served as the key pathway for the O<sub>3</sub> removal only when O<sub>3</sub> pollution  
343 was very serious. The power of O<sub>3</sub> column amount on surface O<sub>3</sub> concentration seemed to be lower  
344 than those of most meteorological factors, suggesting that vertical transport of ambient O<sub>3</sub> was  
345 complex. Although socioeconomic factors and land use types were not dominant factors for the O<sub>3</sub>  
346 pollution in Tibetan Plateau, they still cannot be ignored in the present study because the predictive  
347 performance would worsen if these variables were excluded from the model. It was widely  
348 acknowledged that the emissions of NO<sub>x</sub> and VOCs focused on the developed urban areas with high  
349 population density especially in the remote plateau (Zhang et al., 2007; Zheng et al., 2017).  
350 Compared with the urban land, the grassland played more important role on the O<sub>3</sub> pollution in  
351 Tibetan Plateau. It was thus supposed that the grassland was widely distributed on Tibetan Plateau,



which could release a large amount of biogenic volatile organic compounds (BVOCs) (Fang et al., 2015). It was well known that photochemical reaction of BVOCs and  $\text{NO}_x$  in the presence of sunlight was beneficial to the  $\text{O}_3$  formation (Calfapietra et al., 2013). Furthermore, Fang et al. (2015) confirmed that the BVOC emission in Tibetan Plateau displayed a remarkable increase in the wet seasons.

### 3.3 The spatial distribution of estimated 8-h $\text{O}_3$ concentration over Tibetan Plateau

Figure 6 depicts the spatial distribution of the 8-h  $\text{O}_3$  level estimated by the novel RF-GAM model. The spatial distribution pattern modelled by the RF-GAM model showed the similar characteristic with the result simulated by previous studies except North Tibetan Plateau (Liu et al., 2018). The estimated 8-h  $\text{O}_3$  concentration displayed the highest value in some cities of North Tibetan Plateau such as Huangnan ( $73.48 \pm 4.53 \mu\text{g}/\text{m}^3$ ) and Hainan ( $72.24 \pm 5.34 \mu\text{g}/\text{m}^3$ ), followed by the cities in the central region including Lhasa ( $65.99 \pm 7.24 \mu\text{g}/\text{m}^3$ ) and Shigatse ( $65.15 \pm 6.14 \mu\text{g}/\text{m}^3$ ), and the lowest one in some cities of Southeast Tibetan Plateau such as Aba ( $55.17 \pm 12.77 \mu\text{g}/\text{m}^3$ ). The spatial pattern of 8-h  $\text{O}_3$  concentration is highly consistent with the result predicted by Liu et al. (2018) using CMAQ model, while it is not in agreement with the result estimated by Zhan et al. (2018) using RF-STK model. The difference of the present study and Zhan et al. (2018) focuses on the North Tibetan Plateau, which lacks of monitoring site and remains the higher uncertainty. Firstly, it might be contributed by the weakness of RF-STK mentioned above. Moreover, Zhan et al. (2018) only used the ground-level measured data in 2015 to establish the model and the data in new sites since 2015 were not incorporated into the model, which could increase the model uncertainty (Zhan et al., 2018). As shown in Fig. 6, most of the cities in Qinghai province (e.g., Huangnan, Hainan, and Guoluo) generally showed the higher 8-h  $\text{O}_3$  concentration over Tibetan Plateau, which



374 was in a good agreement with the spatial distribution of  $O_3$  column amount (Fig. S2). Besides, some  
 375 cities in Tibet such as Shigatse and Lhasa also showed the higher 8-h  $O_3$  levels. It was supposed that  
 376 the precursor (e.g.,  $NO_x$  and VOCs) emissions in these regions were significantly higher than those  
 377 in other cities of Tibetan Plateau (Fig. S3). Zhang et al. (2007) used the satellite data to observe that  
 378 the higher VOCs and  $NO_x$  emission focused on the residential area with high population density in  
 379 the remote Tibetan Plateau. Apart from the effect of anthropogenic emission, the meteorological  
 380 conditions could be also the important factors for the 8-h  $O_3$  concentration. As shown in Fig. S4-  
 381 S10, the higher blh and sp in the Northeast Tibetan Plateau were beneficial to the  $O_3$  formation  
 382 through the reaction of VOC and OH radical, leading to the higher 8-h  $O_3$  concentration in these  
 383 cities (Ou et al., 2015). In addition, the lower tp occurred in North Tibetan Plateau and Northeast  
 384 Tibetan Plateau, both of which were unfavorable to the ambient  $O_3$  removal (Yoo et al., 2014). In  
 385 contrast, the higher tp observed in the Southeast Tibetan Plateau resulted in the slight  $O_3$  pollution.

#### 386 3.4 The temporal variation of the simulated 8-h $O_3$ concentration over Tibetan Plateau

387 The annually mean estimated 8-h  $O_3$  concentration in Tibetan Plateau displayed the slow  
 388 increase from  $64.74 \pm 8.30 \mu\text{g}/\text{m}^3$  to  $66.45 \pm 8.67 \mu\text{g}/\text{m}^3$  2005 through 2015, whereas it decreased  
 389 from the peak to  $65.87 \pm 8.52 \mu\text{g}/\text{m}^3$  during 2015-2018 (Fig. 7). Based on the Mann-Kendall method  
 390 (Fig. 8a), it was found that the surface  $O_3$  concentration exhibited the slight increase as the whole,  
 391 while the increase degree was not significant ( $p > 0.05$ ). Besides, it should be noted that the  $O_3$   
 392 concentrations in various regions showed different increase speed. As depicted in Fig. 8b, we found  
 393 that the 8-h  $O_3$  concentrations in North, West, and East Tibetan Plateau displayed significant  
 394 increase trend by the speed of  $1\text{-}3 \mu\text{g}/\text{m}^3$  during 2005-2018. The middle region of Tibetan Plateau  
 395 showed the moderate increase trend by the speed of  $0\text{-}1 \mu\text{g}/\text{m}^3$ . However, the 8-h  $O_3$  concentration



396 in Shigatse and Sannan even displayed the decrease trend 2005 through 2018.

397 Besides, the 8-h  $O_3$  concentration in Tibetan Plateau displayed significantly seasonal

398 discrepancy. The estimated 8-h  $O_3$  level in Tibetan Plateau followed the order of spring ( $75.00 \pm 8.56$

399  $\mu\text{g}/\text{m}^3$ ) > summer ( $71.05 \pm 11.13 \mu\text{g}/\text{m}^3$ ) > winter ( $56.39 \pm 7.42 \mu\text{g}/\text{m}^3$ ) > autumn ( $56.13 \pm 8.27 \mu\text{g}/\text{m}^3$ )

400 (Fig. 9 and Tab. 3). The 8-h  $O_3$  concentrations in most of prefecture-level cities showed the similarly

401 seasonal characteristics with the overall seasonal variation in Tibetan Plateau. Based on the result

402 summarized in Tab. S1, it was found that the key precursors of ambient  $O_3$  (e.g., VOCs,  $\text{NO}_x$ )

403 generally displayed the higher emissions in winter compared with other seasons. However, the

404 seasonal distribution of ambient  $O_3$  concentration was not in accordance with the precursor

405 emissions, suggesting that the meteorological factors might play more important roles on ambient

406  $O_3$  concentration. It was well known that the higher air temperature in spring and summer was

407 closely related to the low sp and high sund, both of which were beneficial to the  $O_3$  formation

408 (Sitnov et al., 2017). Although summer showed the highest air temperature and the longest sunshine

409 duration, the higher rainfall amount in summer decreased the ambient  $O_3$  concentration via wet

410 deposition (Li et al., 2017b; Li et al., 2019b). Moreover, the highest blh occurred in spring, which

411 was favorable to the strong stratosphere-troposphere exchange process in Tibetan Plateau (Skerlak

412 et al., 2014). Therefore, the 8-h  $O_3$  concentration in summer and winter were relatively lower than

413 that in spring. Nonetheless, the 8-h  $O_3$  levels in Diqing, Sannan, and Nyingchi displayed the highest

414 values in spring ( $56.38 \pm 7.87$ ,  $73.90 \pm 5.97$ , and  $73.22 \pm 2.77 \mu\text{g}/\text{m}^3$ ), followed by winter ( $45.88 \pm 7.05$ ,

415  $61.71 \pm 4.32$ , and  $62.24 \pm 3.63 \mu\text{g}/\text{m}^3$ ) and summer ( $44.35 \pm 5.90$ ,  $61.00 \pm 5.86$ , and  $59.60 \pm 2.33 \mu\text{g}/\text{m}^3$ ),

416 and the lowest ones in autumn ( $37.45 \pm 5.76$ ,  $54.70 \pm 3.13$ , and  $53.84 \pm 2.06 \mu\text{g}/\text{m}^3$ ). The lower  $O_3$  level

417 in summer than winter was mainly attributable to the higher precipitation observed in the summer



418 of these cities (Fig. S9). In addition, it should be noted that the  $\text{NO}_x$  and VOCs emissions of South  
419 Tibetan Plateau (e.g., Sannan) exhibited the higher values in winter compared with other seasons.

### 420 3.5 The nonattainment days over Tibetan Plateau during 2005-2018

421 The annually mean nonattainment days in the 19 prefecture-level cities over Tibetan Plateau are  
422 summarized in Tab. 2.  $100 \mu\text{g}/\text{m}^3$  was regarded as the critical value for the 8-h  $\text{O}_3$  level by World  
423 Health Organization (WHO). The nonattainment days denoted total days with the 8-h  $\text{O}_3$   
424 concentration higher than  $100 \mu\text{g}/\text{m}^3$ . Although the annually mean 8-h  $\text{O}_3$  concentrations in all of  
425 the cities over Tibetan Plateau did not exceed the critical value, not all of the regions experienced  
426 excellent air quality in the long period (2005-2018). Some cities of Qinghai province including  
427 Huangnan, Haidong, and Guoluo suffered from 45, 40, and 40 nonattainment days each year (Fig.  
428 10 and Tab. 4). Besides, some cities in the South Tibetan Plateau such as Shigatse and Sannan also  
429 experienced more than 40 nonattainment days each year, suggesting that Tibetan Plateau was still  
430 faced of the risk for  $\text{O}_3$  pollution. Fortunately, some remote cities such as Ali, Ngari, and Qamdo  
431 did not experience the excessive  $\text{O}_3$  pollution all the time, which was ascribed to the low precursor  
432 emissions and appropriate meteorological conditions. It should be noted that the nonattainment days  
433 in the region with high  $\text{O}_3$  concentration showed the significantly seasonal difference, whereas the  
434 seasonal difference was not remarkable in the city with low  $\text{O}_3$  pollution. As shown in Tab. 2, it  
435 should be noted that nearly all of the nonattainment days could be detected in spring and summer,  
436 which was in good agreement with the  $\text{O}_3$  levels in different seasons, indicating that the  $\text{O}_3$  pollution  
437 issue should be paid more attention in spring and summer.

### 438 4. Summary and implication

439 In the present study, we developed a novel hybrid model (RF-GAM) based on multiple



440 explanatory variables to estimate the surface 8-h O<sub>3</sub> concentration across the remote Tibetan Plateau.

441 The 10-fold cross-validation method demonstrated that RF-GAM achieved excellent performance

442 with the highest R<sup>2</sup> value (0.76) and lowest root mean square error (RMSE) (14.41 µg/m<sup>3</sup>) compared

443 with other model including RF-STK, RF, BPNN, XGBoost, GRNN, ElmanNN, and ELM models.

444 Moreover, the unlearning ground-measured O<sub>3</sub> data validated that the RF-GAM model showed the

445 better extrapolation performance (R<sup>2</sup>=0.67, RMSE=25.68 µg/m<sup>3</sup>). The result of variable importance

446 suggested that time, sund, and sp were key factors for the surface 8-h O<sub>3</sub> concentration over Tibetan

447 Plateau. Based on the RF-GAM model, we found that the estimated 8-h O<sub>3</sub> concentration exhibited

448 notably spatial variation with the highest value in some cities of North Tibetan Plateau such as

449 Huangnan (73.48±4.53 µg/m<sup>3</sup>) and Hainan (72.24±5.34 µg/m<sup>3</sup>) and the lowest one in some cities of

450 Southeast Tibetan Plateau such as Aba (55.17±12.77 µg/m<sup>3</sup>). Besides, we also found that the O<sub>3</sub>

451 level displayed a slow increase from 64.74±8.30 µg/m<sup>3</sup> to 66.45±8.67 µg/m<sup>3</sup> 2005 through 2015,

452 while the O<sub>3</sub> concentration decreased to 65.87±8.52 µg/m<sup>3</sup> in 2018. The estimated 8-h O<sub>3</sub> level in

453 Tibetan Plateau showed the significantly seasonal discrepancy with the order of spring (75.00±8.56

454 µg/m<sup>3</sup>) > summer (71.05±11.13 µg/m<sup>3</sup>) > winter (56.39±7.42 µg/m<sup>3</sup>) > autumn (56.13±8.27 µg/m<sup>3</sup>).

455 Based on the critical value set by WHO, most of the cities in Tibetan Plateau shared with the

456 excellent air quality, while several cities (e.g., Huangnan, Haidong, and Guoluo) still suffered from

457 more than 40 nonattainment days each year.

458 The RF-GAM model for O<sub>3</sub> estimation has several limitations. First of all, the O<sub>3</sub> estimation in

459 North Tibetan Plateau might show some uncertainties because the ground-level monitoring site is

460 very scarce, and thus we cannot validate the reliability of predicted value in the region without

461 monitoring site. Secondly, our approach did not include data on emission inventory, or traffic count



because the continuous emissions of NO<sub>x</sub> and VOCs were not open access. At last, we only focused on the temporal variation of surface O<sub>3</sub> concentration in recent ten years, and the short-term O<sub>3</sub> data cannot reflect the response of O<sub>3</sub> pollution to climate change. In the future work, we should combine more explanatory variables such as long-term NO<sub>x</sub> and VOCs emissions to retrieve the surface O<sub>3</sub> level over Tibetan Plateau in the past decades.

#### Author contributions

This study was conceived by Rui Li and Hongbo Fu. Statistical modelling was performed by Rui Li, Yilong Zhao, Ya Meng, Wenhui Zhou and Ziyu Zhang. Rui Li drafted the paper.

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## Figure and table captions

**Fig. 1** The geographical locations and annually mean 8-h  $O_3$  concentrations in the ground-observed sites (red dots) over Tibetan Plateau during 2014-2018. The elevation data are collected from geographical and spatial data cloud at a 30-m spatial resolution.

**Fig. 2** The workflow for predicting the spatiotemporal distributions of 8-h  $O_3$  levels.

**Fig. 3** Density scatterplots of model fitting and cross-validation result at a daily level. (a), (b), (c), (d), (e), (f), (g), and (h) represent RF-GAM, RF-STK, RF, GRNN, XGBoost, BPNN, ElmanNN, and ELM models, respectively. The red dotted line denotes the fitting linear regression line. The full names of MPE and RMSE are mean prediction error ( $\mu g/m^3$ ) and root mean squared prediction error ( $\mu g/m^3$ ), respectively.

**Fig. 4** The transferring ability validation of RF-GAM method based on the measured daily 8-h  $O_3$  concentration during December 2013-May 2014.

**Fig. 5** The variable importance of predictors in the final RF-GAM model.

**Fig. 6** The mean value of estimated 8-h  $O_3$  concentration during 2005-2018 over Tibetan Plateau.

**Fig. 7** The inter-annual variation of predicted 8-h  $O_3$  level ( $\mu g/m^3$ ) from 2005 to 2018 across Tibetan Plateau.

**Fig. 8** The trend analysis of predicted 8-h  $O_3$  concentration. (a) and (b) represent the result of Mann-Kendall method and discrepancy of estimated  $O_3$  level during 2005-2018 across Tibetan Plateau.

**Fig. 9** The seasonal variability of estimated 8-h  $O_3$  level across Tibetan Plateau. (a), (b), (c), and (d) represent the predicted 8-h  $O_3$  concentrations in spring, summer, autumn, and winter, respectively.

**Fig. 10** The spatial distributions of nonattainment days in Tibetan Plateau during 2005-2018.

**Tab. 1** The  $R^2$  values, RMSE, and MPE of RF-GAM in four seasons over Tibetan Plateau.



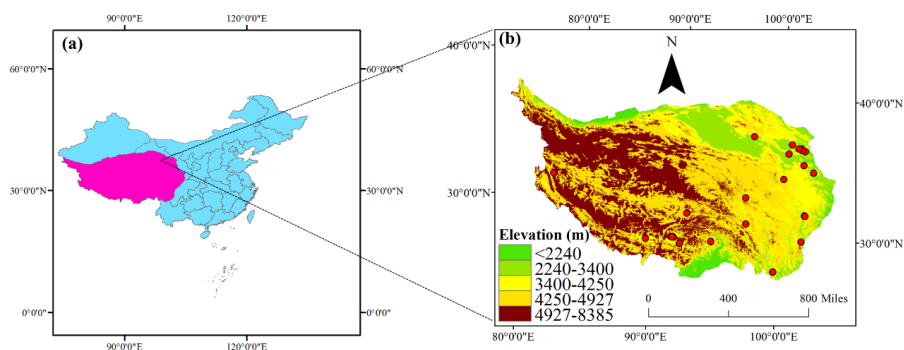
**Tab. 2** The  $R^2$  values, RMSE, and MPE of RF-GAM in different provinces over Tibetan Plateau.

**Tab. 3** The estimated 8-h  $O_3$  concentration in 19 prefecture-level cities over Tibetan Plateau during four seasons including spring, summer, autumn, and winter.

**Tab. 4** The mean nonattainment days (8-h  $O_3$  level  $>100 \mu\text{g}/\text{m}^3$ ) in 19 prefecture-level cities over Tibetan Plateau each year.

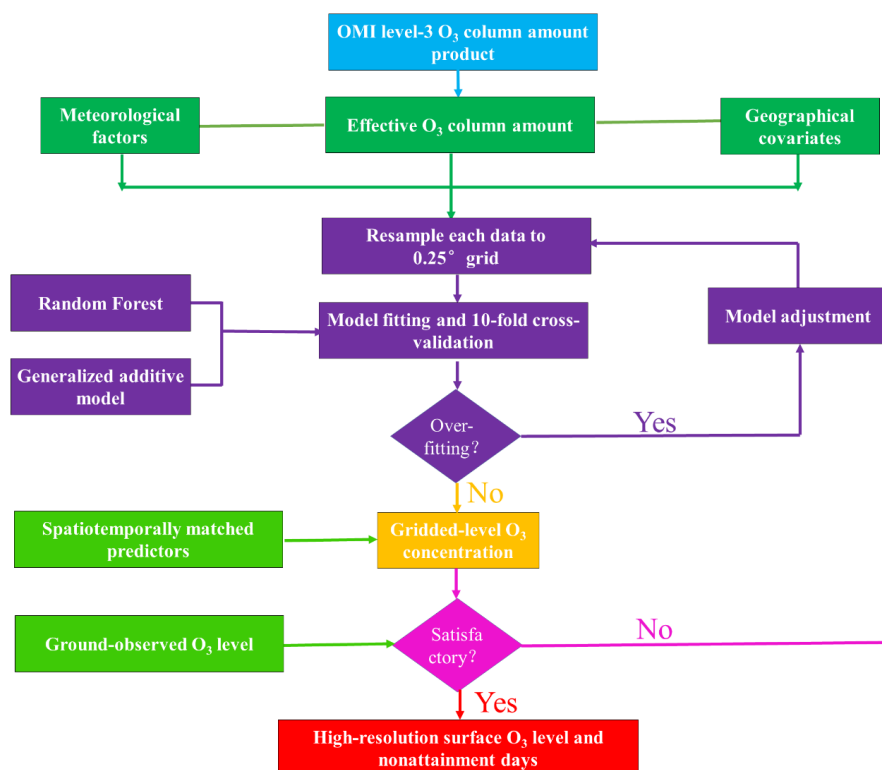


**Fig. 1**



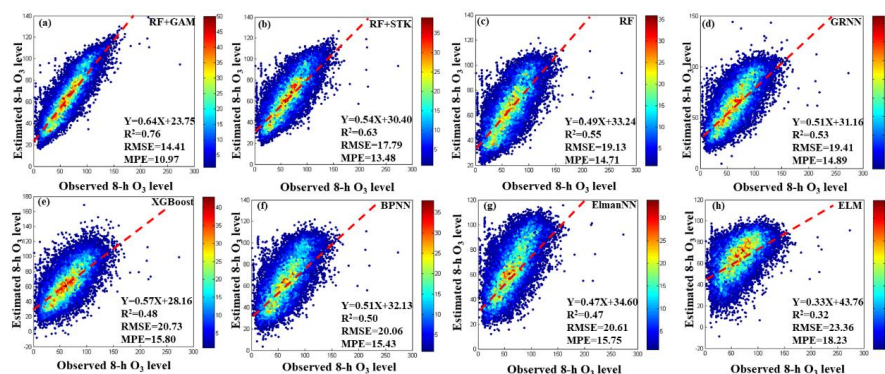


**Fig. 2**



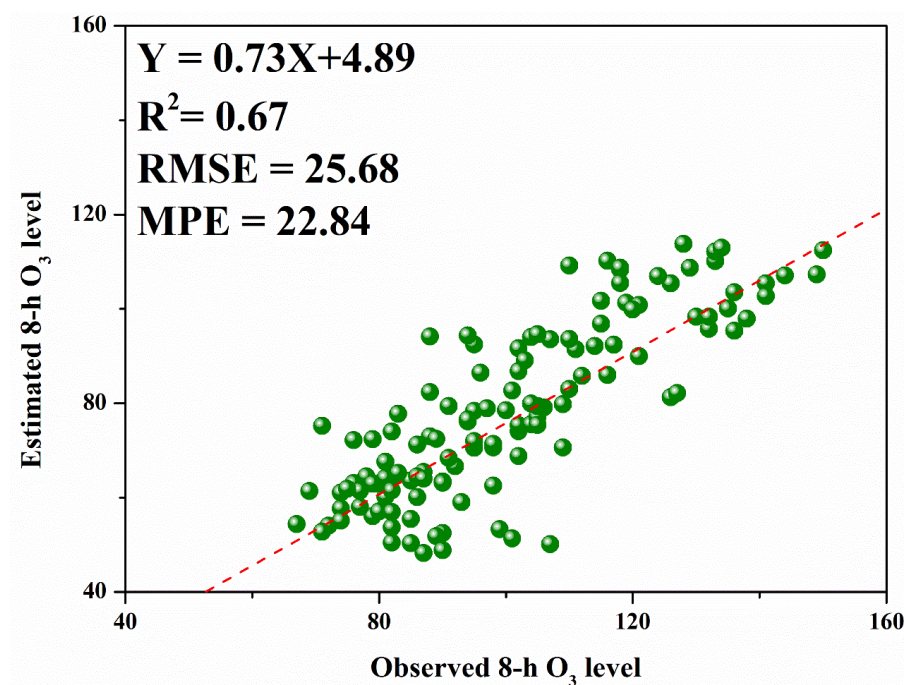


**Fig. 3**



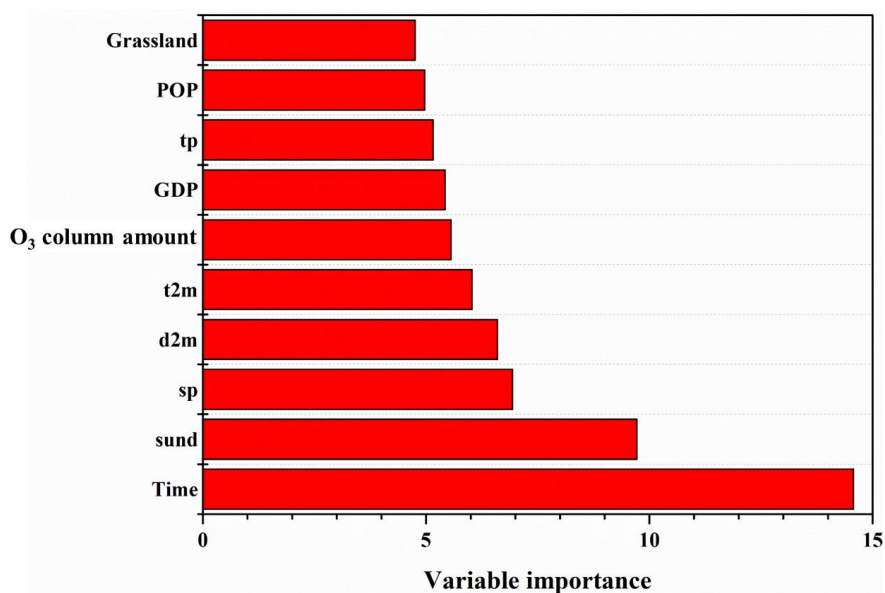


**Fig. 4**



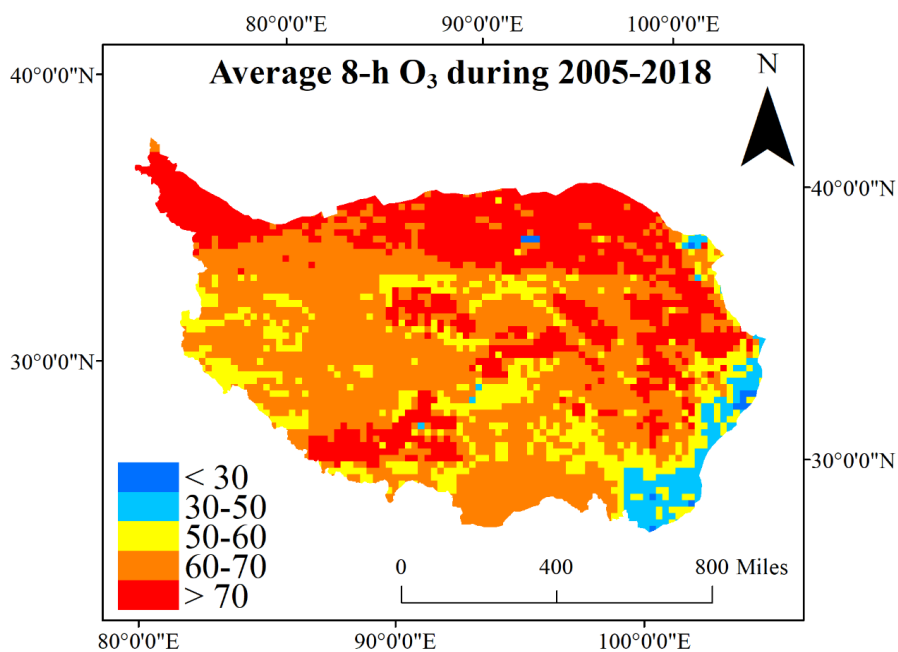


**Fig. 5**



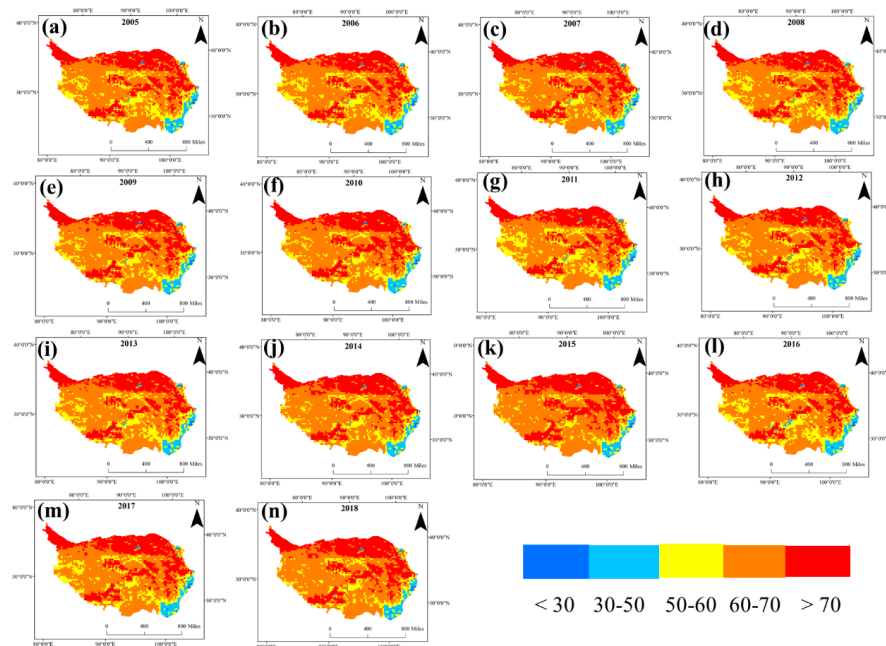


**Fig. 6**



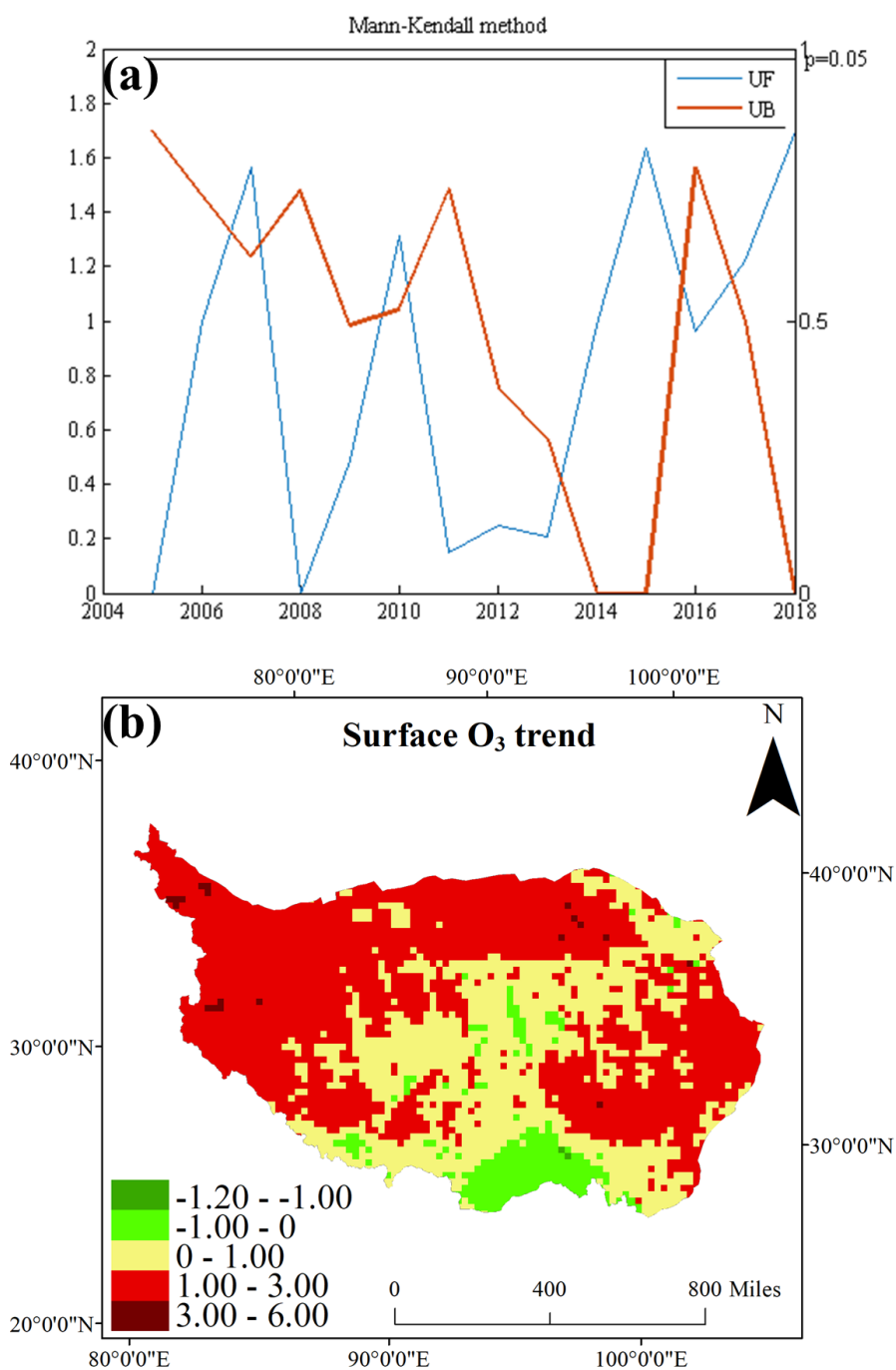


**Fig. 7**



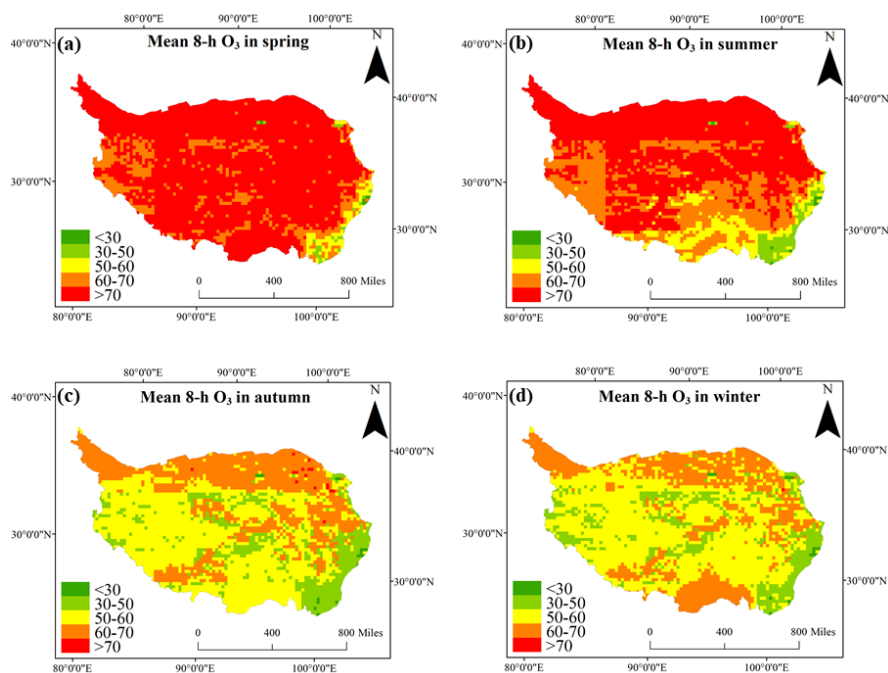


**Fig. 8**



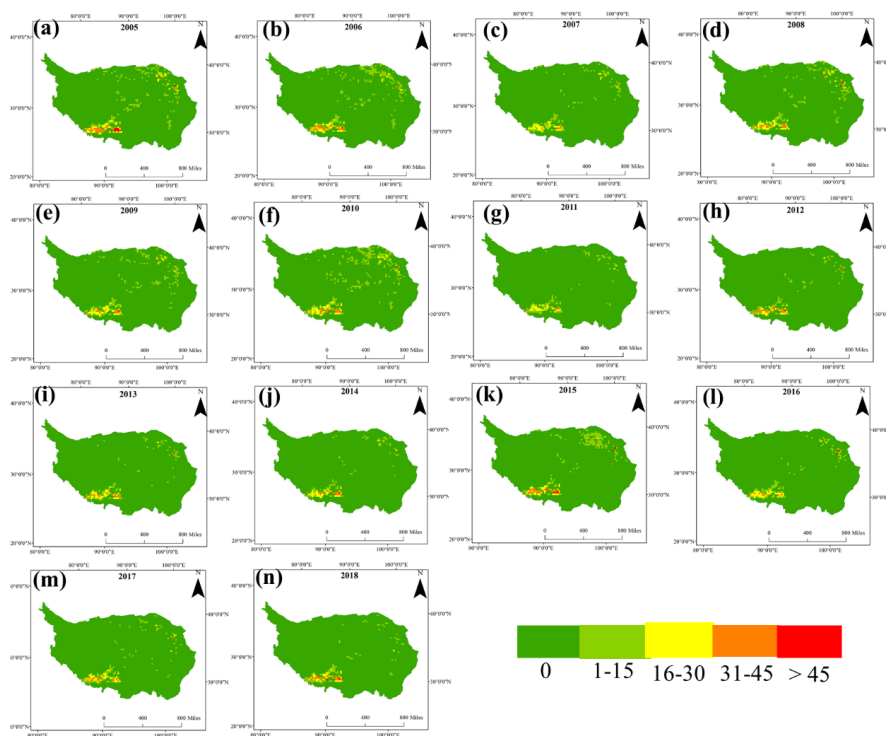


**Fig. 9**





**Fig. 10**





**Tab. 1**

|       | Spring | Summer | Autumn | Winter |
|-------|--------|--------|--------|--------|
| $R^2$ | 0.64   | 0.74   | 0.67   | 0.69   |
| RMSE  | 15.32  | 15.13  | 13.23  | 14.58  |
| MPE   | 11.94  | 11.75  | 10.52  | 11.44  |



**Tab. 2**

|       | Tibet | Qinghai | Gansu | Sichuan | Yunnan |
|-------|-------|---------|-------|---------|--------|
| $R^2$ | 0.69  | 0.70    | 0.74  | 0.71    | 0.54   |
| RMSE  | 14.81 | 14.83   | 13.65 | 13.23   | 12.49  |
| MPE   | 11.24 | 11.33   | 10.88 | 10.08   | 10.20  |



**Tab. 3**

|          | Province | Spring      | Summer      | Autumn      | Winter      | Annual      |
|----------|----------|-------------|-------------|-------------|-------------|-------------|
| Aba      | Sichuan  | 65.61±14.30 | 59.46±14.32 | 45.55±12.03 | 47.95±10.55 | 55.17±12.77 |
| Ngari    | Tibet    | 71.34±3.12  | 70.10±3.57  | 53.14±3.67  | 51.84±3.69  | 62.21±3.34  |
| Qamdo    | Tibet    | 72.52±4.29  | 62.74±5.79  | 52.06±4.01  | 55.42±3.09  | 61.10±3.93  |
| Diqing   | Yunnan   | 56.38±7.87  | 44.35±5.90  | 37.45±5.76  | 45.88±7.05  | 46.22±6.51  |
| Gannan   | Gansu    | 76.77±9.73  | 73.27±10.67 | 54.74±8.33  | 54.72±6.95  | 65.60±8.91  |
| Ganzi    | Sichuan  | 69.38±10.99 | 61.45±11.58 | 48.49±8.79  | 50.94±6.62  | 58.06±9.48  |
| Guoluo   | Qinghai  | 80.12±5.12  | 76.13±5.83  | 58.86±5.71  | 57.38±4.66  | 68.77±5.25  |
| Haibei   | Qinghai  | 78.18±10.21 | 78.84±10.31 | 60.90±9.69  | 57.48±9.78  | 69.47±9.99  |
| Haidong  | Qinghai  | 74.20±10.34 | 73.70±9.12  | 53.61±8.11  | 51.02±9.60  | 63.84±9.21  |
| Hainan   | Qinghai  | 83.01±5.36  | 82.27±5.72  | 61.57±5.39  | 58.96±5.44  | 72.24±5.34  |
| Haixi    | Qinghai  | 79.39±6.88  | 79.48±7.79  | 60.78±7.48  | 57.71±6.99  | 69.99±7.24  |
| Huangnan | Qinghai  | 85.21±4.98  | 83.01±4.66  | 61.95±4.18  | 60.62±4.49  | 73.48±4.53  |
| Lhasa    | Tibet    | 80.08±9.63  | 70.13±8.42  | 55.86±5.78  | 55.85±5.19  | 65.99±7.24  |
| Nagqu    | Tibet    | 74.59±5.13  | 70.46±6.69  | 54.60±5.16  | 53.53±4.83  | 63.83±5.23  |
| Shigatse | Tibet    | 77.31±8.62  | 69.66±7.69  | 55.93±4.58  | 55.57±4.72  | 65.15±6.14  |
| Sannan   | Tibet    | 73.90±5.97  | 61.00±5.86  | 54.70±3.13  | 61.71±4.32  | 63.04±4.00  |
| Xining   | Qinghai  | 77.43±10.27 | 77.84±9.44  | 58.19±9.29  | 54.72±10.04 | 67.77±9.70  |
| Yushu    | Qinghai  | 77.35±5.55  | 73.34±6.37  | 56.12±5.53  | 55.02±5.01  | 66.05±5.50  |
| Nyingchi | Tibet    | 73.22±2.77  | 59.60±2.33  | 53.84±2.06  | 62.24±3.63  | 62.40±2.20  |



**Tab. 4**

|          | Spring | Summer | Autumn | Winter | Annual |
|----------|--------|--------|--------|--------|--------|
| Aba      | 0      | 0      | 0      | 0      | 0      |
| Ngari    | 0      | 0      | 0      | 0      | 0      |
| Qamdo    | 0      | 0      | 0      | 0      | 0      |
| Diqing   | 0      | 0      | 0      | 0      | 0      |
| Gannan   | 0      | 1      | 0      | 0      | 1      |
| Ganzi    | 13     | 2      | 0      | 0      | 15     |
| Guoluo   | 19     | 21     | 0      | 0      | 40     |
| Haibei   | 0      | 0      | 0      | 0      | 0      |
| Haidong  | 22     | 18     | 0      | 0      | 40     |
| Hainan   | 14     | 12     | 1      | 0      | 27     |
| Haixi    | 1      | 1      | 0      | 0      | 2      |
| Huangnan | 23     | 22     | 0      | 0      | 45     |
| Lhasa    | 12     | 7      | 0      | 0      | 19     |
| Nagqu    | 24     | 14     | 0      | 0      | 38     |
| Shigatse | 28     | 13     | 0      | 0      | 41     |
| Sannan   | 33     | 7      | 0      | 0      | 40     |
| Xining   | 2      | 1      | 0      | 0      | 3      |
| Yushu    | 0      | 0      | 0      | 0      | 0      |
| Nyingchi | 0      | 0      | 0      | 0      | 0      |