



Insights on Atmospheric Oxidation Processes by Performing Factor Analyses on Sub-ranges of Mass Spectra

- 3 Yanjun Zhang¹, Otso Peräkylä¹, Chao Yan¹, Liine Heikkinen¹, Mikko Äijälä¹, Kaspar R. Daellenbach¹,
- 4 Qiaozhi Zha¹, Matthieu Riva^{1,2}, Olga Garmash¹, Heikki Junninen^{1,3}, Pentti Paatero¹, Douglas Worsnop^{1,4}, and
- 5 Mikael Ehn¹
- 6 ¹ Institute for Atmospheric and Earth System Research / Physics, Faculty of Science, University of Helsinki,
- 7 Helsinki, 00014, Finland
- 8 ² Univ Lyon, Université Claude Bernard Lyon 1, CNRS, IRCELYON, F-69626, Villeurbanne, France
- ⁹ Institute of Physics, University of Tartu, Tartu, 50090, Estonia
- 10 ⁴ Aerodyne Research, Inc., Billerica, MA 01821, USA
- 11
- 12 Corresponding author: yanjun.zhang@helsinki.fi
- 13

14 Abstract

15 With the recent developments in mass spectrometry, combined with the strengths of factor analysis 16 techniques, our understanding of atmospheric oxidation chemistry has improved significantly. The 17 typical approach for using techniques like positive matrix factorization (PMF) is to input all measured 18 data for the factorization in order to separate contributions from different sources and/or processes to 19 the total measured signal. However, while this is a valid approach for assigning the total signal to 20 factors, we have identified several cases where useful information can be lost if solely using this 21 approach. For example, gaseous molecules emitted from the same source can show different temporal 22 behaviors due differing loss terms, like condensation at different rates due to different molecular 23 masses. This conflicts with one of PMF's basic assumptions of constant factor profiles. In addition, 24some ranges of a mass spectrum may contain useful information, despite contributing only minimal 25 fraction to the total signal, in which case they are unlikely to have a significant impact on the 26 factorization result. Finally, certain mass ranges may contain molecules formed via pathways not 27 available to molecules in other mass ranges, e.g. dimeric species versus monomeric species. In this 28 study, we attempted to address these challenges by dividing mass spectra into sub-ranges and 29 applying the newly developed binPMF method to these ranges separately. We utilized a dataset from 30 a chemical ionization atmospheric pressure interface time-of-flight (CI-APi-TOF) mass spectrometer 31 as an example. We compare the results from these three different ranges, each corresponding to 32 molecules of different volatilities, with binPMF results from the combined range. Separate analysis 33 showed clear benefits in dividing factors for molecules of different volatilities more accurately, in





resolving different chemical processes from different ranges, and in giving a chance for highmolecular-weight molecules with low signal intensities to be used to distinguish dimeric species with different formation pathways. In addition, daytime dimer formation (diurnal peak around noon) was identified, which may contribute to NPF in Hyytiälä. Also, dimers from NO₃ oxidation were separated by the sub-range binPMF, which would not be identified otherwise. We recommend PMF users to try running their analyses on selected sub-ranges in order to further explore their datasets.

41 **1 Introduction**

42 Huge amounts of volatile organic compounds (VOC) are emitted to the atmosphere every year 43 (Guenther et al., 1995;Lamarque et al., 2010), which play a significant role in atmospheric chemistry 44 and affect the oxidative ability of the atmosphere. The oxidation products of VOC can contribute to 45 the formation and growth of secondary organic aerosols (Kulmala et al., 2013;Ehn et al., 2014;Kirkby 46 et al., 2016; Troestl et al., 2016), affecting air quality, human health, and climate radiative forcing 47 (Pope III et al., 2009;Stocker et al., 2013;Zhang et al., 2016;Shiraiwa et al., 2017). Thanks to the 48 advancement in mass spectrometric applications, like the aerosol mass spectrometer (AMS) 49 (Canagaratna et al., 2007) and chemical ionization mass spectrometry (CIMS) (Bertram et al., 50 2011; Jokinen et al., 2012; Lee et al., 2014) our capability to detect these oxidized products, as well as 51 our understanding of the complicated atmospheric oxidation pathways in which they take part, have 52 been greatly enhanced.

Monoterpenes (C10H16), one common group of VOC emitted in forested areas, have been shown to 53 54be a large source of atmospheric secondary organic aerosol (SOA). The oxidation of monoterpenes 55 produces a wealth of different oxidation products (Oxygenated VOC, OVOC), including highly 56 oxygenated organic molecules (HOM) with molar yields in the range of a few percent, depending on 57 the specific monoterpene and oxidant (Ehn et al., 2014; Bianchi et al., 2019). Bianchi et al. (2019) 58 summarized that HOM can be either Extremely Low Volatility Organic Compounds (ELVOC), Low 59 Volatility Organic Compounds (LVOC), or Semi-volatile Organic Compounds (SVOC) 60 (classifications by Donahue et al. 2012), depending on their exact structures. For less oxygenated 61 products, the majority are likely to fall into the SVOC or the Intermediate VOC (IVOC) range. The 62 volatility of the OVOC will determine their dynamics, including their ability to contribute to the 63 formation of SOA and new particles (Bianchi et al., 2019;Buchholz et al., 2019). 64 The recent developments of CIMS techniques has allowed researchers to observe unprecedented

numbers of OVOC, in real-time (Riva et al., 2019). This ability to measure thousands of compounds

66 is a great benefit, but also a large challenge for the data analyst. For this reason, factor analytical

67 techniques have often been applied to reduce the complexity of the data by finding co-varying signals





68 that can be grouped into common factors (Huang et al., 1999). For aerosol and gas-phase mass 69 spectrometry, positive matrix factorization, PMF (Paatero and Tapper, 1994;Zhang et al., 2011) has 70 been the most utilized tool. The factors have then been attributed to sources (e.g. biomass burning 71organic aerosol) or processes (e.g. monoterpene ozonolysis) depending on the application and ability 72 to identify spectral signatures (Yan et al., 2016; Zhang et al., 2017). In the vast majority of these PMF 73 applications to mass spectra, the mass range of ions has been maximized in order to provide as much 74 input as possible for the algorithm. This approach was certainly motivated in early application of 75 PMF on e.g. offline filters, with chemical information of metals, water-soluble ions, and organic and 76 elemental carbon (OC and EC), where the number of variables is counted in tens, and the number of 77 samples in tens or hundreds (Zhang et al., 2017). However, with gas-phase CIMS, we often have up 78 to a thousand variables, with hundreds or even thousands of samples, meaning that the amount of data 79 itself is unlikely to be a limitation for PMF calculation. In this work, we aimed to explore potential 80 benefits of dividing the spectra into sub-ranges before applying factorization analysis.

81 An inherent requirement of factorization approaches is that the factor profiles, in this case the relative 82 abundancies of ions in the mass spectra, of each factor stay nearly constant. Due to the complexity 83 and number of atmospheric processes affecting the formation, transformation, and loss of VOC, 84 OVOC and aerosol, this often does not hold, and is one of the main limitations of factorization 85 approaches. Given the different volatilities of OVOC, it may even be expected that molecules from 86 the same source may have very different loss time scales, which may affect the factor analysis. 87 Volatility issue has been studied and reported for AMS data, with different volatilities of various OA 88 types (Huffman et al., 2009;Crippa et al., 2014;Paciga et al., 2016;Äijälä et al., 2017). Semi-volatile 89 oxygenated organic aerosol (SV-OOA) and Low-volatility oxygenated organic aerosol (LV-OOA) 90 can both be mainly produced from biogenic sources, but get separated based on different volatilities 91 by PMF (El Haddad et al., 2013). Sekimoto et al. (2018) found that the two profiles resolved with 92 VOC emitted from biomass burning had different estimated volatilities. As the volatility of a molecule 93 is linked to its molecular mass (Peräkylä et al., 2019), it may be beneficial to apply PMF separately 94 to mass ranges where one can expect the loss processes to be similar, thereby resulting in more 95 constant factor profiles. In this way, distinct sources are hopefully separated by PMF, with minimized 96 influence of differing volatilities from one source.

97 The number of PMF or other factorization studies utilizing CIMS data remains very limited. 98 "Traditional" PMF analyses have so far, to our knowledge, only been applied to nitrate-based 99 chemical ionization atmospheric pressure interface time-of-flight (CI-APi-TOF) data (Yan et al., 100 2016;Massoli et al., 2018). One study has also utilized non-negative matrix factorization (NNMF) to 101 look at diurnal trends of Iodide ToF-CIMS data (Lee et al., 2018). The lack of more studies utilizing





PMF, or other factorization techniques, on CIMS data is likely partly due to the complexity of the 102 103 data, with multiple overlapping ions hampering HR peak fitting (Zhang et al., 2019). In addition, 104 variable factor profiles may hamper PMF's ability to correctly separate the factors. The two CI-APi-105 TOF studies utilized the nearly the entire measured spectrum (from around 200 Th to around 600 Th), 106 either in unit mass resolution (UMR) or high resolution (HR) peak fitting data (Yan et al., 107 2016; Massoli et al., 2018). Massoli et al. (2018) estimated the volatility of the molecules they detected, 108 finding that all the six extracted factors had notable contributions from IVOC, SVOC and (E)LVOC. 109 These compound groups will have clearly different loss mechanisms, and thereby loss rates, 110 suggesting that variation in factor profiles is inevitable, even if the source was identical for all 111 molecules in the factor. We hypothesize that this effect further hampers the correct factorization, and 112 further that this effect can be reduced by dividing the spectra into separate ranges, with each sub-113 range containing molecules with roughly similar loss mechanisms and rates.

114 As an additional motivation to separate different ranges from the mass spectrum, it is not only the 115 loss mechanisms, but also the formation pathways that may differ. For example, atmospheric 116 oxidation chemistry of organics is, to a large extent, the chemistry of peroxy radicals (RO₂) (Orlando 117 and Tyndall, 2012). These RO₂ are initiated by VOC reacting with oxidants like ozone, or the hydroxyl 118 (OH) or nitrate (NO₃) radicals, while their termination occurs mainly by bimolecular reactions with 119 NO, HO₂ and/or other RO₂. Some product molecules can be formed from many of the three 120 termination pathways, while for example ROOR "dimers" can only be formed from RO₂+RO₂ 121 reactions (Berndt et al., 2018a;Berndt et al., 2018b). This also means that there can be six different 122 pathways to form dimers from the same precursors VOC, by combining RO₂ formed from the same 123 or different oxidants. As an example of the latter, an ROOR dimer can contain one moiety produced 124 from ozone oxidation and another moiety from NO_3 oxidation (Yan et al., 2016). Thus, their 125 concentration is dependent on both the precursor VOC concentration, and the concentrations of both 126 oxidants. Such a molecule will not have a direct equivalent in any of the monomer products, 127 dependent on only one oxidant, which again may complicate the separation of such factors by PMF, 128 if the entire spectrum is analyzed once. However, if separating the monomer and dimer products

129 before PMF analysis, separation of different formation pathways can potentially become simpler.

130 Recently, we proposed a new PMF approach, binPMF, to simplify the analysis of mass spectral data

131 (Zhang et al., 2019). This method divides the mass spectrum into narrow bins, typically some tens of

132 bins per integer mass, depending on the mass resolving power of the instrument, before performing

133 PMF analyses. In this way, binPMF does not require any time-consuming, and potentially subjective

134 high resolution peak fitting, and can thus be utilized for data exploration at a very early stage of data

135 analysis. Data preparation is nearly as simple as in the case of UMR analysis, yet it utilizes much





- more spectral information as it does not sum up signal over all ions at each integer mass. In addition to saving time and effort in data analysis, the results are less sensitive to mass calibration fluctuations. Finally, the binning also greatly increases the number of input variables, which has the advantage that factor analysis with smaller mass ranges becomes more feasible, as more meaningful variation is supplied to the algorithm.
- 141 We designed this study to explore the benefits of separate analysis of different mass ranges from mass
- spectra. We used a previously published ambient dataset measured by a CI-APi-TOF, and conducted
- 143 binPMF analysis with three different mass ranges, roughly corresponding to different volatility ranges.
- 144 We compared the results from the sub-range analyses with each other and with results from binPMF
- 145 run on the combined ranges. We found that more meaningful factors are separated from our dataset
- by utilizing the sub-ranges, and believe that this study will provide new perspectives for future studies
- 147 analyzing gas-phase CIMS data.
- 148

149 2 Methodology

The focus of this work is on retrieving new information from mass spectra by applying new analytical approaches. Therefore, we chose a dataset that has been presented earlier, though without PMF analysis, by Zha et al. (2018), and was also used in the first study describing the binPMF method (Zhang et al., 2019). The measurements are described in more details below in section 2.1, while the data analysis techniques used in this work are presented in section 2.2.

- 155 2.1 Measurements
- 156 2.1.1 Ambient site
- 157 The ambient measurements were conducted at the Station for Measuring Ecosystem–Atmosphere
- 158 Relations (SMEAR) II in Finland (Hari and Kulmala, 2005) as part of the Influence of Biosphere-
- 159 Atmosphere Interactions on the Reactive Nitrogen budget (IBAIRN) campaign (Zha et al, 2018).
- 160 Located in the boreal environment in Hyytiälä, SMEAR II is surrounded with coniferous forest and
- 161 has limited anthropogenic emission sources nearby. Diverse measurements of meteorology, aerosol
- 162 and gas phase properties are continuously conducted at the station. Details about the meteorological
- 163 conditions and temporal variations of trace gases during IBAIRN campaign are presented by Zha et
- 164 al. (2018) and Liebmann et al. (2018).
- 165 2.1.2 Instrument and data
- 166 Data were collected with a nitrate (NO₃)-based chemical ionization atmospheric pressure interface
- 167 time-of-flight mass spectrometer (CI-APi-TOF, Jokinen et al., 2012) with about 4000 Th Th⁻¹ mass
- resolving power, at ground level in September, 2016. In our study, the mass spectra were averaged to
- 169 1 h time resolution from September 6^{th} to 22^{nd} for further analysis. We use the thomson (Th) as the





170 unit for mass/charge, with 1 Th = 1 Da/e, where e is the elementary charge. As all the data discussed 171in this work are based on negative ion mass spectrometry, we will use the absolute value of the 172mass/charge, although the charge of each ion will be negative. The masses discussed in this work 173 includes the contribution from the nitrate ion, 62, unless specifically mentioned. Furthermore, as the 174technique is based on soft ionization with NO_3^- ions, any multiple charging effects are unlikely, and 175therefore the reported mass/charge values in thomson can be considered equivalent to the mass of the 176 ion in Da. 177 The forest site of Hyytiälä is dominated by monoterpene emissions (Hakola et al., 2006). The main 178 feature of previous CI-APi-TOF measurements in Hyytiälä (Ehn et al., 2014; Yan et al., 2016) has 179 been a bimodal distributions of HOM, termed monomers and dimers, as they are formed of either one 180 or two RO₂ radicals, respectively. For the analysis in this study, we chose three mass/charge (m/z)181 ranges of 50 Th each (Figure 1), corresponding to regions between which we expect differences in 182 formation or loss mechanisms. In addition to regions with HOM monomers and HOM dimers, one 183 range was chosen at lower masses, in a region presumably mainly consisting of molecules that are 184 less likely to condense onto aerosol particles (Peräkylä et al., 2019). 185 2.2 Positive matrix factorization (PMF) 186 After the model of PMF was developed (Paatero and Tapper, 1994), numerous applications have been 187 conducted with different types of environmental data (Song et al., 2007;Ulbrich et al., 2009;Yan et 188 al., 2016; Zhang et al., 2017). By reducing dimensionality of the measured dataset, PMF model greatly 189 simplifies the data analysis process with no requirement for prior knowledge of sources or pathways

as essential input. The main factors can be further interpreted with their unique/dominant markers(elements or masses).

The basic assumption for PMF modelling is mass balance, which assumes that ambient concentration
of a chemical component is the sum of contributions from several sources or processes, as shown in
equation (1).

195

$\mathbf{X} = \mathbf{TS} \times \mathbf{MS} + \mathbf{R} \tag{1}$

196 In equation (1), **X** stands for the time series of measured concentration of different variables (m/z in 197 our case), **TS** represents the temporal variation of factor contributions, **MS** stands for factor profiles 198 (mass spectral profiles), and **R** is the residual as the difference of the modelled and the observed data. 199 The matrices **TS** and **MS** are iteratively calculated by a least-squares algorithm utilizing uncertainty 200 estimates, to pursue minimized Q value as shown in equation (2), where S_{ij} is the estimated 201 uncertainty, an essential input in PMF model.

202
$$Q = \sum \sum \left(\frac{R_{ij}}{S_{ij}}\right)^2$$
(2)





203	PMF model was conducted by multi-linear engine (ME-2) (Paatero, 1999) interfaced with Source
204	Finder (SoFi, v6.3) (Canonaco et al., 2013). Signal-to-noise ratio (SNR) was calculated as SNR <i>ij</i> =
205	abs (Xij) / abs (Sij) . When the Signal-to-noise ratio (SNR) is below 1, the signal of X_{ij} will be down-
206	weighted by replacing the corresponding uncertainty S_{ij} by S_{ij} /SNR _{ij} (Visser et al., 2015). Future
207	studies should pay attention to the potential risk when utilizing this method since down-weighting
208	low signals element-wise will create a positive bias to the data. Robust mode was operated in the
209	PMF modelling, where outliers $\left(\frac{R_{ij}}{S_{ij}}\right > 4$) were significantly down-weighted (Paatero, 1997).
210	2.3 binPMF
211	As a newly developed application of PMF for mass spectral data, binPMF has no requirement for
212	chemical composition information, while still taking advantage of the HR mass spectra, saving effort
213	and time (Zhang et al., 2019). To explore the benefits of analyzing separated mass ranges, we applied
214	binPMF to the three separated ranges. The three ranges were also later combined for binPMF analysis
215	as comparison with the previous results. The PMF model requires both data matrix and error matrix
216	as input, and details of the preparation of data and error matrices are described below.
217	2.3.1 Data matrix
218	Different from normal UMR or HR peak fitting, in binPMF, the mass spectra are divided into small
219	bins after baseline subtraction and mass axis calibration. Linear interpolation was first conducted to
220	the mass spectra with a mass interval of 0.001 Th. Then the interpolated data was averaged into bins
221	of 0.02 Th width. We selected three ranges for further analysis based on earlier studies (Ehn et al.,
222	2014;Yan et al., 2016;Bianchi et al., 2019;Peräkylä et al., 2019).
223	- Range 1, m/z 250 – 300 Th, 51 unit masses × 25 bins per unit mass = 1275 bins/variables,
224	consisting mainly of molecules with five to nine carbon atoms and four to nine oxygen atoms
225	in our dataset.
226	- Range 2, m/z 300 – 350 Th, $51 \times 25 = 1275$ bins, mainly corresponding to HOM monomer
227	products, featured with nine to ten C- and seven to ten O-atoms.
228	- Range 3, $m/z 510 - 560$ Th, $51 \times 30 = 1530$ bins, mainly corresponding to HOM dimer products,
229	with carbon numbers of sixteen to twenty and eleven to fifteen O-atoms.
230	For a nominal mass N , the signal region included in further analyses was between N -0.2 Th and N +0.3
231	Th for Range 1 and 2, and between N -0.2 Th and N +0.4 Th for Range 3. The data were averaged into
232	1-h time resolution and in total we had 384 time points in the data matrix.

233 2.3.2 Error matrix





- The error matrix represents the estimated uncertainty for each element of the data matrix and is crucial for iterative calculation of the Q minimum. Equation (3) is used for error estimation (Polissar et al., 1998).
- 237

240

253

$$S_{ii} = \sigma_{ii} + \sigma_{\text{noise}} \tag{3}$$

where S_{ij} represents the uncertainty of m/z j at time *i*, σ_{ij} stands for counting statistics uncertainty and is estimated as follows,

$$\sigma_{ij} = a \times \frac{\sqrt{I_{ij}}}{\sqrt{t_s}} \tag{4}$$

where *I* is the signal intensity term, in unit of counts per second (cps), t_s stands for length of averaging in seconds, while *a* is an empirical coefficient to compensate for unaccounted uncertainties (Allan et al., 2003;Yan et al., 2016) and is 1.28 in our study as previously estimated from laboratory experiments (Yan et al., 2016). The σ_{noise} term was estimated as the median of the standard deviations from signals in the bins in the region between nominal masses, where no physically meaningful signals are expected.

247 3 Results

248 3.1 General overview of the dataset/spectrum

During the campaign, in autumn, 2016, the weather was overall sunny and humid with average temperature of 10.8 °C and relative humidity (RH) of 87% (Zha et al., 2019). The average concentration of NO_x and O₃ was 0.4 ppbv and 21 ppbv, respectively. The average total HOM concentration was ~ 10^8 molecules cm⁻³.







- Figure 1. Example of mass spectrum with 1-h time resolution measured from a boreal forest
- 255 environment during the IBAIRN campaign (at 18:00, Finnish local time, UTC+2). The mass
- spectrum was divided into three parts and three sub-ranges were chosen from different parts for
- 257 further analysis in our study. The nitrate ion (62 Th) is included in the mass.
- Figure 1 shows the 1 h averaged mass spectrum taken at 18:00 on September 12, as an example of
- the analyzed dataset. In addition to exploring the benefits of this type of sub-range analysis in relation
- to different formation or loss pathways, separating into sub-ranges may also aid factor identification
- for low-signal regions. As shown in Figure 1, there is a difference of 1-2 orders of magnitude in the
- signal intensity between Range 3 and Ranges 1-2. If all Ranges are run together, we would expect
- that the higher signals from Ranges 1 and 2 will drive the factorization. While if run separately,
- separating formation pathways of dimers in Range 3 will likely be easier. As dimers have been shown
- to be crucial for the formation of new aerosol particles from monoterpene oxidation (Kirkby et al.,
- 266 2016;Troestl et al., 2016;Lehtipalo et al., 2018), this information may even be the most critical in 267 some cases, despite the low contribution of these peaks to the total measured signal.
- binPMF was separately applied to Range 1, 2, 3, and a 'Range combined' which comprised all the three sub-ranges. All the PMF runs for the four ranges were conducted from two to ten factors and repeated three times for each factor number, to assure the consistency of the results. Factorization results and evolution with increasing factor number are briefly described in the following sections, separately for each Range (sections 3.2 - 3.5). More detailed discussion and comparison between the
- 273 results are presented in Section 4.
- 274 3.2 binPMF on Range 1 (250 300 Th)
- As has become routine (Zhang et al., 2011;Craven et al., 2012), we first examined the mathematical parameters of our solutions. From two to ten factors, Q/Q_{exp} decreased from 2.8 to 0.7 (Fig S1 in supplementary information), and after three factors, the decreasing trend was gradually slowing down and approaching one, which is the ideal value for Q/Q_{exp} as a diagnostic parameter. The unexplained variation showed a decline from 18% to 8% from two to ten factors.
- In the two-factor results, two daytime factors were separated, with peak time both at 14:00 15:00.
 One factor was characterized by large signals at 250 Th, 255 Th, 264 Th, 281 Th, 283 Th, 295 Th,
 297 Th. The other factor was characterized by large signals at 294 Th, 250 Th, 252 Th, 264 Th, 266
 Th, 268 Th, and 297 Th. In Hyytiälä, as reported in previous studies, odd masses observed by the
 nitrate CI-APi-TOF are generally linked to monoterpene-derived organonitrates during the day (Ehn
 et al., 2014;Yan et al., 2016). When the factor number increased to three, the two earlier daytime
- 286 factors remained similar with the previous result, while a new factor appeared with a distinct sawtooth
- shape in the diurnal cycle. The main marker in the spectral profile was 276 Th, with a clear negative





mass defect. When one more factor was added, the previous three factors remained similar as in the three-factor solution, and a new morning factor was resolved, with 264 Th and 297 Th dominant in the mass spectral profile, and a diurnal peak at 11:00.

291 As the factor number was increased, more daytime factors were separated, with similar spectral 292 profiles to existing daytime factors and various peak times. No nighttime factors were found in the 293 analysis even when the factor number reached ten. We chose the four-factor result for further 294 discussion, and Figure 2 shows the result of Range 1, with spectral profile, time series, diurnal cycle 295 and factor contribution. As shown in Figure 2d, Factors 1-3 are all daytime factors, while Factor 4 296 has a sawtooth shape, which is caused by contamination, mainly by perfluorinated acids, of the inlet's 297 automated zeroing every three hours during the measurements (Zhang et al., 2019). The zeroing 298 periods have been removed from the dataset before binPMF analysis, but the contamination factor 299 was still resolved. This factor is discussed in more detail in sections 4.1.1 and 4.1.4.







303 3.3 binPMF on Range 2 (300-350 Th)

This range covers the monoterpene HOM monomer range, and binPMF results have already been discussed by Zhang et al. (2019) as a first example of the application of binPMF on ambient data. Our input data here is slightly different. In the previous study, the 10 min automatic zeroing every three hours was not removed before averaging to 1 hour time resolution while here, we have removed this data. Overall, the results are similar as in our earlier study, and therefore the result are just briefly





- summarized below for further comparison and discussion in Section 4. Similar to Range 1, both the Q/Q_{exp} (2.2 to 0.6) and unexplained variation (16% to 8%) declined with the increased factor number
- 311 from two to ten.

312 When the factor number was two, one daytime and one nighttime factor were separated, with diurnal 313 peak times at 14:00 and 17:00, respectively. The nighttime factor was characterized by masses at 340 314 Th, 308 Th and 325 Th (monoterpene ozonolysis HOM monomers (Ehn et al., 2014)) and remained 315 stable throughout the factor evolution from two to ten factors. With the addition of more factors, no 316 more nighttime factors got separated while the daytime factor was further separated and more daytime 317 factors appeared, peaking at various times in the morning (10:00 am), at noon or in the early afternoon 318 (around 14:00 pm and 15:00 pm). High contribution of 339 Th can be found in all the daytime factor 319 profiles. As the factor number reached six, a contamination factor appeared, characterized by large 320 signals at 339 Th and 324 Th, showing negative mass defects (Figure S2 in the Supplement). The 321 factor profile is nearly identical to the contamination factor determined in Zhang et al. (2019), where 322 the zeroing periods were not removed, causing larger signals for the contaminants. In our dataset, 323 where the zeroing periods were removed, no sawtooth pattern was discernible in the diurnal trend, 324 yet it could still be separated even though it only contributed 3% to Range 2. More about the 325 contamination factors from different ranges will be discussed in Section 4.1.4. Since the aim of this 326 study is mainly to explore the benefits of analyzing different ranges of the mass spectrum, we chose 327 to show the four-factor result below, to simplify the later discussion and comparison. Figure 3 shows 328 four-factor result of Range 2, with spectral profile, time series, diurnal cycle and factor contribution.







- 330 Figure 3 Four-factor result for Range 2, for (a) factor spectral profiles, (b) factor contribution, (c)
- time series and (d) diurnal trend. Details on the factors' naming schemes are shown in Table 1.
- 332 3.4 binPMF on Range 3 (510-560 Th)
- Range 3 represents mainly the monoterpene HOM dimers (Ehn et al., 2014). Similar to Range 1 and
- 334 2, both the Q/Q_{exp} (1.5 to 0.6) and unexplained variation (18% to 15%) showed decreasing trend with
- the increased factor number (2-10). As can be seen from Figure 1, data in Range 3 had much lower
- signals, compared to that of the Range 1 and 2, explaining the higher unexplained variation for Range337
- 338 In the two-factor result for Range 3, one daytime and one nighttime factor appeared, with diurnal 339 peak times at noon and 18:00, respectively. The nighttime factor was characterized by ions at 510 Th, 340 524 Th, 526 Th, 542 Th, and 555 Th, 556 Th, while the daytime factor showed no dominant marker 341 masses, yet with relatively high signals at 516 Th, 518 Th and 520 Th. As the number of factors 342 increased to three, one factor with almost flat diurnal trend was separated, with dominant masses of 343 510 Th, 529 Th, 558 Th. Most peaks in this factor had negative mass defects, and this factor was 344 again linked to a contamination factor. The four-factor result resolved another nighttime factor with 345 a dominant peak at 555 Th, and effectively zero contribution during daytime. As the factor number 346 was further increased, the new factors seemed like splits from previous factors with similar spectral 347 profiles. We therefore chose four-factor result also for Range 3 (results shown in Fig. 4) for further 348 discussion.





Figure 4 Four-factor result for Range 3, for (a) factor spectral profiles, (b) factor contribution, (c) time series and (d) diurnal trend. Details on the factors' naming schemes are shown in Table 1.





- 352 3.5 binPMF on Range Combined (250-350 Th & 510-560 Th)
- 353 As comparison to the previous three ranges, we conducted the binPMF analysis on Range Combined,
- 354 which is the combination of the three ranges. The results of this range are fairly similar to those of
- Ranges 1 and 2, as could be expected since the signal intensities in these ranges were much higher
- 356 than in Range 3. As the number of factors increased (2-10), both the Q/Q_{exp} (1.3 to 0.6) and
- unexplained variation (16% to 8%) showed a decreasing trend.
- 358 In the two-factor result, one daytime factor and one nighttime factor were separated. In the nighttime 359 factor, most masses were found at even masses, and the fraction of masses in Range 3 was much 360 higher than that in daytime factor. In contrast, in the daytime factor, most masses were observed at 361 odd masses and the fraction of signal in Range 3 was much lower. During the day, photochemical 362 reactions increase the concentration of NO, which serves as peroxy radical (RO₂) terminator and often 363 outcompetes RO_2 cross reactions in which dimers can be formed (Ehn et al., 2014). Thus, the 364 production of dimers is suppressed during the day, yielding instead a larger fraction of organic nitrates, 365 as has been shown also previously (Yan et al., 2016). 366 With the increase of the number of factors, more daytime factors were resolved with different peak
- times. When the factor number reached seven, a clear sawtooth-shape diurnal cycle occurred, i.e. the contamination factor, caused by the zeroing. As more factors were added, no further nighttime factors were separated, and only more daytime factors appeared. To simplify the discussion and inter-range comparison, we also here chose the four-factor result for further analysis, as it already provided enough information for our main goal in this study. Figure 5 shows the four-factor result of Range Combined, with spectral profile, time series, diurnal cycle and factor contribution. The signals in range of 510-560 Th were enlarged 100-fold to be visible.







374

375 Figure 5 Four-factor result for Range Combined, for (a) factor spectral profiles, (b) factor

contribution, (c) time series and (d) diurnal trend. Details on the factors' naming schemes are shown

377 in Table 1.

378 4 Discussion

379 In Section 3, results by binPMF analysis were shown for Ranges 1, 2, 3 and Combined. In this section, 380 we discuss and compare the results from the different ranges. To simplify the inter-range comparison, 381 we chose four-factor results for all the four ranges, with the abbreviations shown in Table 1. From 382 Range 1, three daytime factors and a contaminations factor were separated. In Range 2, three daytime 383 factors and one nighttime factor (abbreviated as R2F4_N) were resolved. The R2F4_N factor was characterized by signals at 308 Th (C10H14O7·NO3-), 325 Th (C10H15O8·NO3-), and 340 Th 384 385 $(C_{10}H_{14}O_9 \cdot NO_3)$, and can be confirmed as monoterpene ozonolysis products (Ehn et al., 2014; Yan et 386 al., 2016). With the increase of factor number to six, the contamination factor got separated also in 387 this mass range. In Range 3, one daytime factor, two nighttime factors and a contamination factor 388 were separated. The first nighttime factor (R3F2 N1) had large peaks at 510 Th ($C_{20}H_{32}O_{11}$ ·NO₃⁻) 389 and 556 Th ($C_{20}H_{30}O_{14}$ ·NO₃⁻), dimer products that have been identified during chamber studies of 390 monoterpene ozonolysis (Ehn et al., 2014). The molecule observed at 510 Th has 32 H-atoms, 391 suggesting that one of the RO₂ involved would have been initiated by OH, which is formed during 392 the ozonolysis of alkenes such as monoterpenes at nighttime (Atkinson et al., 1992;Paulson and





393	Orlando, 1996). The other nighttime factor (R3F3_N2) was dominated by ions at 523 Th
394	$(C_{20}H_{31}O_8NO_3 \cdot NO_3^-)$ and 555 Th $(C_{20}H_{31}O_{10}NO_3 \cdot NO_3^-)$, representing nighttime monoterpene
395	oxidation involving NO $_3$. As these dimers contain only one N-atom, and 31 H-atoms, we can assume
396	that they are formed from reactions between an RO_2 formed from NO_3 oxidation and another RO_2
397	formed by ozone oxidation. These results match well with the profiles in a previous study by Yan et
398	al. (2016). The results of Range Combined are very similar to Range 2, with one nighttime factor and
399	three daytime factors. The contamination factor was separated with increase of factor number to seven.
400	

401

Table 1. Summary of PMF results for the different mass ranges

Range	Factor number	Factor name ^a	Dominant peaks	Peak time
	1	R1F1 D1	250, 255, 295, 297	15:00
	2	R1F2 D2	250, 252, 294	15:00
1 (250 - 300 Th)	3	R1F3 D3	264.297	11:00
	4	R1F4 C	276	_b
	1		307, 309, 323, 325, 339,	15:00
A (200 - 250 TH)	2	R2F2_D2	310, 326, 339,	14:00
2 (300 - 350 Th)	3	R2F3_D3	339	11:00
	4	R2F4_N	308, 325, 340	18:00
	1	R3F1_D	516, 518, 520, 528, 540	12:00
2 (510 5(0.771))	2	R3F2_N1	510, 524, 542, 556	18:00
3 (510 – 560 1 h)	3	R3F3_N2	523, 555	22:00
	4	R3F4_C	510, 558	_ ^b
	1	RCF1_D1	250, 255, 295, 339	15:00
	2	RCF2_D2	250, 252, 294, 339	14:00
Combined $(1, 2, 3)$	3	RCF3_D3	264, 297, 339	11:00
	4	RCF4 N	308, 340, 510, 524, 555, 556	18:00

402 ^a Factor name is defined with range name, factor number and name. For example, RxFy represents Factor y in Range x.

403 RC stands for Range Combined. For the factor name, D is short for daytime, N for Nighttime, C for contamination.

⁴⁰⁴ ^b The contamination factor in Range 1 shows sawtooth pattern; while in Range 3 shows no diurnal pattern.

405

406 4.1 Comparison of different ranges

407 4.1.1 Time series correlation

408 In Figure 6, the upper panels show the time series correlations among the first three ranges. As 409 expected based on the results above, generally the daytime factors, and the two nighttime 410 monoterpene ozonolysis factors (R2F4_N and R3F2_N1) correlated well, respectively. However, the 411 contamination factors did now show strong correlation between different ranges, even though 412 undoubtedly from the same source. More about the contamination factors will be discussed in Section 413 4.1.4. The lower panels in Figure 6 displays the correlations between the first three ranges and the 414 Range Combined, and clearly demonstrates that the results of Range Combined is mainly controlled 415 by high signals from Range 1 and 2. More detailed aspects of the comparison between factors in



418



- 416 different ranges is given in the following sections. The good agreements between factors from
- 417 different subranges also help to verify the robustness of the solutions.



Figure 6 Time series correlations among Range 1, 2, 3 (upper panels a-c), and between the first three ranges and the Range Combined (lower panels d-f). The abbreviations for different factors are the same in Table 1, with F for factor, D for daytime, N for nighttime and C for contamination, e.g. F1D1 for Factor 1 daytime 1. The coefficient of determination, R², is marked in each subplot by a number shown in the right upper corners and by the blue colors, with stronger blue indicating higher R².

424 4.1.2 Daytime factor comparison

As mentioned above, with increasing number of factors, usually more daytime factors will be resolved, reflecting the complicated daytime photochemistry. The three daytime factors between Range 1 and 2 agreed with each other quite well (Figure 6a). However, R1F1_D1 and R2F1_D1 did not show strong correlation with the only daytime factor in Range 3 (R3F1_D), while the other two daytime factors in both Range 1 and 2, i.e. R1F2_D2, R1F3_D3, R2F2_D2, R2F3_D3, correlated well with R3F1_D from Range 3.

The 1st daytime factors from Range 1 and 2, R1F1_D1 and R2F1_D1, were mainly characterized by odd masses 255 Th, 281 Th, 283 Th, 295 Th, 297 Th, 307 Th, 309 Th, 311 Th, 323 Th, 325 Th, 339 Th. The factors are dominated by organonitrates. Organic nitrate formation during daytime is generally associated with the termination of RO₂ radicals by NO. This termination step is mutually exclusive with the termination of RO₂ with other RO₂, leading to dimer formation. If the NO concentration is the limiting factor for the formation of these factors, the low correlations between the NO-terminated monomer factors, and the dimer factors, is to be expected. In contrast, if the other





- daytime factors mainly depend on oxidant and monoterpene concentrations, some correlation
 between those, and the daytime dimer factor, is to be expected, as shown in Figure 6b, c.
- 440 All the spectral profiles resolved from Range Combined binPMF analysis inevitably contained mass
- 441 contribution from 510 560 Th, even the daytime factor from Range Combined (RCF1 D1) which
- 442 did not show clear correlation with R3F1_D from Range 3 (Figure 6e).
- 443 The 2nd and 3rd daytime factors in Range 1 and 2, R1F2_D2, R1F3_D3, R2F2_D2, R2F3_D3, had 444 high correlations with R3F1 D in Range 3. Daytime factors in Range Combined (RCF2 D2 and 445 RCF3 D3) also showed good correlation with R3F1 D in Range 3. However, if we compare R3F1 D 446 and the mass range of 510 - 560 Th of the daytime factors in Range Combined, just with a quick look, 447 we can readily see the difference. The daytime factor separated in Range 3 (R3F1_D) basically has 448 no obvious markers in the profile, and as mentioned above, up to ten factors, there would only be 449 more factors fragmented from the previous factor, with similar spectral profiles, but showed different 450 profile pattern with 510 – 560 Th in RCF2_D2 in Range Combined. The factorization of Range 451 Combined was mainly controlled by Range 1 and 2 due to high signals, and the signals in Range 3 452 are forced to be distributed according to the time series determined by Ranges 1 and 2. Ultimately, 453 this will lead to failure in factor separation for this low-signal range.
- 454 4.1.3 Nighttime factor comparison

Since high-mass dimers are more likely to form at night due to photochemical production of NO in daytime, which inhibits RO₂ + RO₂ reactions, Range 3 had the highest fraction of nighttime signals of all the sub-ranges. While Range 3 produced two nighttime factors, Ranges 2 and Combined showed one, and Range 1 had no nighttime factor. The difference between the two results also indicates the advantage of analyzing monomers and dimers separately.

- 460 The two nighttime factors in Range 3 can be clearly identified as arising from ozonolysis (R3F2 N1) 461 and a mix of ozonolysis and NO_3 oxidation (R3F2 N2) based on the mass spectral profiles, as 462 described above. The organonitrate at 555 Th, $C_{20}H_{31}O_{10}NO_3 \cdot NO_3^-$, is a typical marker for NO₃ 463 radical initiated monoterpene chemistry (Yan et al., 2016). However, several interesting features 464 become evident when comparing to the results of Range 2 and Combined. Firstly, only one nighttime 465 factor (R2F4_N, RCF4_N) was separated in each of these ranges, and that shows clear resemblance 466 with ozonolysis of monoterpenes as measured in numerous studies, e.g. (Ehn et al., 2012;Ehn et al., 467 2014). Secondly, the high correlation found in Figure 6b between the ozonolysis factors (i.e.,
- 468 R2F4_N, R3F2_N1, RCF4_N), further supports the assignment. However, this factor is the only
- 469 nighttime factor in the monomer range, suggesting that NO₃ radical chemistry of monoterpenes in
- 470 Hyytiälä does not form substantial amounts of HOM monomers. The only way for the CI-APi-TOF
- 471 to detect products of monoterpene-NO₃ radical chemistry may thus be through the dimers, where one





highly oxygenated RO₂ radical from ozonolysis reacts with a less oxygenated RO₂ radical from NO₃

473 oxidation.

474 In the results by Yan et al. (2016) the combined UMR-PMF of monomers and dimers did yield a considerable amount of compounds in the monomer range also for the NO₃ radical chemistry factor. 475 476 There may be several reasons for this discrepancy. One major cause for differences between the spring 477 dataset of Yan et al. (2016) and the autumn dataset presented here, is that nighttime concentrations 478 of HOM was greatly reduced during our autumn campaign. The cause may have been fairly frequent 479 fog formation during nights, and also the concentration of e.g. ozone decreased nearly to zero during 480 several nights (Zha et al., 2018). It is also possible that the NO₃ radical-related factor by Yan et al. 481 (2016) is probably a mixture of NO₃ and O₃ radical chemistry, while the monomer may thus be 482 attributed to the O_3 part. Alternatively, the different conditions during the two measurement periods, 483 as well as seasonal difference in monoterpene mixtures (Hakola et al., 2012), caused variations in the 484 oxidation pathways.

485 4.1.4 Contamination factor

486 During the campaign, an automated instrument zeroing every three hours was conducted, by 487 switching a valve to pass the air through a HEPA filter. Each zeroing process lasted for 10 min. While 488 the zeroing successfully removed the low-volatile HOM and H₂SO₄, the zeroing process introduced 489 contaminants into the inlet lines. The contaminants were primarily different types of perfluorinated 490 organic acids, often off-gassing from e.g. Teflon tubing. For IVOC contaminants, these would be 491 flushed through the inlet, while (E)LVOC would condense onto the inlet walls and not come off. 492 However, SVOC contaminants may stick to the inlet tubing and slowly evaporate back into the 493 sampled air. We removed all the 10-min zeroing periods, and averaged the data to 1-h time resolution, 494 but contaminants were still identified in all ranges by binPMF.

495 Contamination contributed 10%, 3%, 19% and 4% to Range 1, 2, 3, and Combined, respectively, in 496 the binPMF solutions where the contamination factor was first separated. This also explains why the 497 contamination factor was separated much earlier in Ranges 1 and 3 than in Range 2. However, despite contributing slightly more to Range Combined than to Range 2, the contamination factor was 498 499 separated when the factor number was increased by one in Range Combined. Here, the difference in 500 volatility of the contaminants in the different sub-ranges may play a role, such that the contaminants 501 in different sub-ranges behave differently. Thus, the behavior of the contamination factor across the 502 combined range is not consistent. Therefore, we examined the zeroing effect with finer time 503 resolution, i.e. 1 min, with three of the largest fluorinated compounds in each range of our mass 504 spectrum, (CF₂)₃CO₂HF·NO₃⁻ (275.9748 Th), (CF₂)₅C₂O₄H⁻ (338.9721 Th), and (CF₂)₆CO₂HF·NO₃⁻ 505 (425.9653 Th). Since the overall signal levels were very low for these compounds, the time series





- 506 became very noisy with such high time resolution. This made it impossible to perform HR fitting for 507 the data, and instead we summed up the signal from the mass ranges where we expected unperturbed
- 508 signal from these ions.
- 509 The time series with sawtooth pattern of the three fluorinated compounds is shown in Figure S3 in
- 510 Supplement. From the time series, we selected a period of around three days of the 3-h cycles (25 in
- 511 total), and in Figure 7 the cycles were aligned and superimposed on top of one another, normalized
- 512 by the maximum during the zeroing. The normalized signals of the three compounds are shown in
- 513 light colors, and the mean values shown in **bold** sold lines. This data includes also the zeroing periods
- 514 to highlight the effect, but these periods were removed from the data used for our PMF analyses.



515

Figure 7 Normalized signals for three fluorinated compounds during a 3-h cycle (180 minutes), with $(CF_2)_3CO_2HF\cdot NO_3^-$ (275.9748 Th) in red, $(CF_2)_5C_2O_4H^-$ (338.9721 Th) in blue, and $(CF_2)_6CO_2HF\cdot NO_3^-$ (425.9653 Th) in black. We selected 25 cycles and normalized all the cycles by their individual maximum. The yellow window shows the zeroing time, for around 10 minutes, which has been removed from the data analysis. Light colors display the individual cycles, and the bold solid colors stand for the average for each compounds.

The signals of the three fluorinated compounds increased by 10 to 20 times during the zeros, due to off-gassing either in the filter or in the tubing in the zeroing setup. Immediately after the zeroing was stopped, signals of all three compounds dropped by about 60-90%, followed by a gradual decay. The decay period coincided with our ambient sampling, and therefore these signals are part of our dataset. It is evident that the three fluorinated compounds were from the same source (zeroing process), but due to their different volatilities, they were lost at different rates. This, in turn, means that the spectral signature of this source will change as a function of time, at odds with one of the basic assumptions





of PMF. Panels a and b in Figure S4 displays the temporal correlation with and without zeroing period with 1 min time resolution. The correlation coefficients dropped greatly when the zero period was removed, from 0.9 to 0.3 for R^2 between 276 Th and 339 Th, and 0.8 to 0.1 between 276 Th and 426

532 Th. Similar effect is also found with the 1 h averaged data (Fig. S4c, d).

- 533 This detailed analysis of fluorinated contamination in our system was here merely used as an example 534to show that volatility can impact source profiles over time. In this case, the contamination factor was 535 still identified both from the separate sub-ranges and from the combined data set using binPMF. 536 However, the contamination profile in the combined range is now averaged, compared to that from 537 separate ranges: the fractional contributions of contamination compounds to this profile, vary during 538 the process of each zeroing due to different volatility properties. In Figure S5, contamination factor 539 profiles from Range 3 and Range Combined were compared. It can be clearly seen that the profile of 540 Range Combined is more noisy than that of Range 3, probably due to the varied fractional 541 contributions of contamination compounds to the profile. In ambient data, products from different 542 sources can have undergone atmospheric processing, altering the product distribution. Our aim with 543 this analysis was to highlight the importance of differences in the sink terms due to different 544 volatilities of the products. This may be an important issue for gas phase mass spectrometry analysis, 545 potentially underestimated by many PMF users, as it is likely only a minor issue for aerosol data, for 546 which PMF has been applied much more routinely. If failing to achieve physically meaningful factors 547 using PMF on gas phase mass spectra, our recommendation is to try applying PMF to sub-ranges of the spectrum, where IVOC, SVOC and (E)LVOC could be analyzed separately. 548
- 549 4.2 Atmospheric insights

550 While the previous section discussed several findings with atmospheric implications, we highlight

- 551 two results below, which are particularly intriguing. We also include the correlation matrix of all
- 552 PMF and factors and selected meteorological parameters in Table 2.
- 553

Table 2 Correlation between factors and meteorological parameters and gases

	R1F1 _D1	R1F1 _D2	R1F1 _D3	R1F1 _C	R2F1 _D1	R2F2 _D2	R2F3 _D3	R2F4 _N	R3F1 _D	R3F2 _N1	R3F3 _N2	R3F4 _C	RCF1 _D1	RCF2 _D2	RCF3 _D3	RCF4 _N
O ₃	0.51	0.59	0.35	-0.18	0.47	0.57	0.36	0.43	0.55	0.33	0.27	0.22	0.49	0.57	0.33	0.34
NO	0.13	-0.01	0.24	-0.03	0.18	-0.02	0.24	-0.22	0.13	-0.19	-0.17	0.03	0.13	0.00	0.26	-0.18
NOx	-0.05	-0.22	-0.10	0.09	-0.01	-0.23	-0.11	-0.13	-0.16	-0.21	-0.04	0.04	-0.04	-0.22	-0.09	-0.11
RH	-0.46	-0.80	-0.63	0.30	-0.43	-0.82	-0.64	-0.27	-0.78	-0.39	-0.07	-0.07	-0.43	-0.82	-0.60	-0.21
Т	0.66	0.72	0.40	-0.24	0.65	0.66	0.41	0.39	0.65	0.30	0.14	0.19	0.66	0.68	0.38	0.24
UVB	0.52	0.63	0.82	-0.40	0.52	0.68	0.84	-0.30	0.79	-0.08	-0.27	0.08	0.49	0.68	0.83	-0.29

554

555





556 4.2.1 Daytime dimer formation

557 Dimers are primarily produced during nighttime, due to NO suppressing $RO_2 + RO_2$ reactions in 558 daytime (Ehn et al., 2014; Yan et al., 2016). However, in this study, we found one clear daytime factor 559 in Range 3 (R3F1_D, peak at local time 12:00, UTC+2) by sub-range analysis. With high loadings 560 from even masses including 516, 518, 520, 528, 540 Th, this only daytime factor in dimer range 561 correlated very well with two daytime factors in Ranges 1 and 2 (R1F2_D2, R1F3_D3, R2F2_D2, 562 R2F3 D3) (Figure 6b and c). Strong correlation between R3F1 D with solar radiation was found, 563 with R = 0.79 (Table 2). This may indicate involvement of OH oxidation in producing this factor. 564 As previous studies have shown, dimers greatly facilitate new particle formation (NPF) (Kirkby et 565 al., 2016;Troestl et al., 2016;Lehtipalo et al., 2018), and this daytime dimer factor may represent a 566 source of dimers that would impact the initial stages of NPF in Hyytiälä. Mohr et al. (2017) reported 567 a clear diel pattern of dimers (sum of about 60 dimeric compounds of C₁₆₋₂₀H₁₃₋₃₃O₆₋₉) during NPF 568 events in 2013 in Hyytiälä, with minimum at night and maximum after noon, and estimated these 569 dimers can contribute ~5% of the mass of sub-60 nm particles. The link between the dimers presented 570 in that paper and those reported here will require further studies, as will the proper quantification of 571 the dimer factor identified here. 572 4.2.2 Dimers initiated by NO₃ radicals 573 Previous studies show that NO₃ oxidation of α -pinene, the most abundant monoterpene in Hyytiälä

574 (Hakola et al., 2012), produces fairly little SOA mass (yields <4 %), while β -pinene shows yields of 575 up to 53 % (Bonn and Moorgat, 2002;Nah et al., 2016). The NO₃+ β -pinene reaction results in low 576volatile organic nitrate compounds with carboxylic acid, alcohol, and peroxide functional groups (Fry 577 et al., 2014;Boyd et al., 2015), while NO₃+ α -pinene reaction will typically lose the nitrate functional 578 group and form oxidation products with high vapor pressures (Spittler et al., 2006;Perraud et al., 579 2010). Most monoterpene-derived HOM, including monomers, are low-volatile (Peräkylä et al., 580 2019) and thus a low SOA yield indicates a low HOM yield. Thus, while there are to our knowledge 581 no laboratory studies on HOM formation from NO₃ oxidation of α -pinene, a low yield can be expected 582 based on SOA studies.

583 As discussed in section 4.1.3, a dimer factor (R3F2_N2) was identified as a mix of ozonolysis and

584 NO₃ oxidation processes, dominated by the organonitrate at 555 Th, C₂₀H₃₁O₁₀NO₃·NO₃⁻. However,

unlike the pure ozonolysis dimer factor which had a corresponding monomer factor (R = 0.86 between

- 586 factor R2F4_N and R3 F2_N1), this NO₃-related dimer factor did not have an equivalent monomer
- 587 factor. This suggests that the NO₃ oxidation of the monoterpene mixture in Hyytiälä does not by itself
- 588 form much HOM, but in the presence of RO₂ from ozonolysis, the RO₂ from NO₃ oxidation can take
- 589 part in HOM dimer formation. This further implies that, different from previous knowledge based on





590 single-oxidant experiments in chambers, NO₃ oxidation may have a larger impact on SOA formation 591 in the atmosphere where different oxidants exist concurrently. This highlights the need for future 592 laboratory studies to consider systems with multiple oxidants during monoterpene oxidation 593 experiments, to truly understand the role and contribution of different oxidants, and NO₃ in particular. 594

595 5 Conclusions

596 The recent development in mass spectrometry has greatly improved the detection of atmospheric 597 vapors and their oxidation products. Factor analysis, such as PMF, can reduce dimensionality of the 598 big datasets and extract factors relating to different atmospheric pathways/sources. Optimally, PMF 599 can link laboratory-generated spectra with ambient observations, significantly improving our 600 understanding of complicated atmospheric processes. However, one of PMF's basic assumptions is 601 that factor profiles remain constant in time, yet for atmospheric gas-phase species, varying sources, 602 reactions and sinks may violate this assumption. Some of these variations are likely not addressable 603 in the data analysis stage, but others may be. For example, molecules formed from the same source 604 can have different temporal behaviors due to varying volatilities, and thus condense at different rates. 605 Performing PMF separately with smaller ranges may circumvent this problem. By utilizing the newly 606 presented binPMF approach, more variables can be extracted from a narrow mass range compared to 607 traditional UMR PMF, while preserving more information in the spectrum.

We conducted separate binPMF analysis on three different sub-ranges to explore the potential benefits of such an approach for producing more physically meaningful factors. We utilized ambient data measured by CI-APi-TOF in a boreal forest environment, and selected sub-ranges from the mass spectrum that roughly corresponded to regions where we would expect the molecules to have similar volatilities and formation pathways. Selected ranges were Range 1 (250 – 300 Th), Range 2 (300 – 350 Th), and Range 3 (510 – 560 Th). binPMF was separately applied to these ranges, as well as to the combination of all three for comparison.

615 The different sub-ranges produced some similar and some different factors. First of all, volatility of 616 species indeed affect the PMF results. We could clearly prove the benefit of sub-range binPMF using 617 the contamination factor as an example. We found that different compounds emitted from the same 618 source showed different temporal trends, likely due to differences in volatilities. This increased the 619 difficulties for PMF to separate this source in the combined data set, and the resolved profile was still 620 less accurate than in the analysis for the sub-ranges. We recommend that future studies of gas-phase 621 mass spectra should pay attention to this volatility effect on factor analysis. 622 Secondly, chemistry or sources contributing to the particular range can be better separated. Only the

623 binPMF analysis on Range 3, where HOM dimers are typically observed, resolved two nighttime





- factors, characterized by monoterpene oxidation related to NO₃ and O₃ oxidation. The monoterpene ozonolysis factor was separated from both Range 2 and 3, showing very good correlation between the ranges and mutually verifying the results.
- Thirdly, peaks with smaller signal intensities can be correctly assigned. The signal intensities between different parts of the mass spectrum may vary by orders of magnitude. In the analysis of the combined range, the results were almost completely controlled by the higher signals from Range 1 and 2. The
- 630 separate analysis on Range 3 allowed the low signals to provide important information, such as the
- 631 NO₃ oxidation process. In addition, running binPMF on different separate mass ranges also allows us
- to compare the factors obtained from the different ranges and help to verify the results.
- 633 In addition, daytime dimer formation was identified, presumably initiated by OH/O_3 with a diurnal 634 peak at around noon, which may contribute to NPF in Hyytiälä. Also, based on the sub-range binPMF 635 analysis, we successfully separated NO_3 -related dimers which did not have an equivalent monomer 636 factor. The NO_3 related factor was consistent with earlier observations (Yan et al., 2016), but would 637 not have been identified from this dataset without utilizing the different sub-ranges. In future 638 laboratory experiments, more complex oxidation systems may be useful in order to understand the 639 role NO_3 oxidation in SOA formation.
- In summary, we identified several reasons to recommend PMF users to try running their analysis on selected sub-ranges in addition to the whole spectra. Ultimately, the approach should be study-goal dependent. In some cases the researcher wants a quick factor analysis to explore different features of their data, while in others more accurate and quantitative separation of different sources and atmospheric processes are needed. As binPMF and UMR-PMF both require very little data preparation, we expect that it in most cases will be worth the time for the analyst to test how PMF results look for a few selected sub-ranges of their mass spectra.
- 647
- **Data availability.** The data used in this study are available from the first author upon request: please contact Yanjun Zhang (yanjun.zhang@helsinki.fi).
- Author contributions. ME and YZ designed the study. QZ and MR collected the data; data analysis
- 663 and manuscript writing were done by YZ. All coauthors discussed the results and commented the
- 664 manuscript.
- 665 Competing interests. The authors declare that they have no conflict of interest
- 666 Acknowledgements. We thank the tofTools team for providing tools for mass spectrometry data
- analysis. The personnel of the Hyytiälä forestry field station are acknowledged for help during field
- 668 measurements.





669 Financial support. This research was supported by the European Research Council (Grant 638703-

- 670 COALA), the Academy of Finland (grants 317380 and 320094), and the Vilho, Yrjö and Kalle
- 671 Väisälä Foundation. K.R.D. acknowledges support by the Swiss National Science postdoc mobility
- 672 grant P2EZP2_181599.

673

674 **Reference**

- Äijälä, M., Heikkinen, L., Fröhlich, R., Canonaco, F., Prévôt, A. S. H., Junninen, H., Petäjä, T., Kulmala, M.,
 Worsnop, D., and Ehn, M.: Resolving anthropogenic aerosol pollution types deconvolution and
 exploratory classification of pollution events, Atmos. Chem. Phys., 17, 3165-3197, 10.5194/acp-17-31652017, 2017.
- Allan, J. D., Jimenez, J. L., Williams, P. I., Alfarra, M. R., Bower, K. N., Jayne, J. T., Coe, H., and Worsnop,
 D. R.: Quantitative sampling using an Aerodyne aerosol mass spectrometer 1. Techniques of data
 interpretation and error analysis, Journal of Geophysical Research: Atmospheres, 108, 2003.
- Atkinson, R., Aschmann, S. M., Arey, J., and Shorees, B.: Formation of OH radicals in the gas phase reactions
 of O3 with a series of terpenes, 97, 6065-6073, 10.1029/92jd00062, 1992.
- Berndt, T., Mentler, B., Scholz, W., Fischer, L., Herrmann, H., Kulmala, M., and Hansel, A.: Accretion Product
 Formation from Ozonolysis and OH Radical Reaction of α-Pinene: Mechanistic Insight and the Influence
 of Isoprene and Ethylene, Environmental Science & Technology, 52, 11069-11077,
 10.1021/acs.est.8b02210, 2018a.
- Berndt, T., Scholz, W., Mentler, B., Fischer, L., Herrmann, H., Kulmala, M., and Hansel, A.: Accretion Product
 Formation from Self- and Cross-Reactions of RO2 Radicals in the Atmosphere, Angewandte Chemie
 International Edition in English 57, 3820-3824, 10.1002/anie.201710989, 2018b.
- Bertram, T. H., Kimmel, J. R., Crisp, T. A., Ryder, O. S., Yatavelli, R. L. N., Thornton, J. A., Cubison, M. J.,
 Gonin, M., and Worsnop, D. R.: A field-deployable, chemical ionization time-of-flight mass spectrometer,
 Atmospheric Measurement Techniques, 4, 1471-1479, 10.5194/amt-4-1471-2011, 2011.
- Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crounse, J. D., Wennberg,
 P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop, D. R., Thornton, J. A.,
 Donahue, N., Kjaergaard, H. G., and Ehn, M.: Highly Oxygenated Organic Molecules (HOM) from GasPhase Autoxidation Involving Peroxy Radicals: A Key Contributor to Atmospheric Aerosol, Chemical
 Reviews, 119, 3472-3509, 10.1021/acs.chemrev.8b00395, 2019.
- Bonn, B., and Moorgat, G. K.: New particle formation during a- and b-pinene oxidation by O3, OH and NO3,
 and the influence of water vapour: particle size distribution studies, Atmos. Chem. Phys., 2, 183-196,
 10.5194/acp-2-183-2002, 2002.
- Boyd, C. M., Sanchez, J., Xu, L., Eugene, A. J., Nah, T., Tuet, W. Y., Guzman, M. I., and Ng, N. L.: Secondary
 organic aerosol formation from the β-pinene+NO3 system: effect of humidity and peroxy radical fate,
 Atmos. Chem. Phys., 15, 7497-7522, 10.5194/acp-15-7497-2015, 2015.
- Buchholz, A., Lambe, A. T., Ylisirniö, A., Li, Z., Tikkanen, O. P., Faiola, C., Kari, E., Hao, L., Luoma, O.,
 Huang, W., Mohr, C., Worsnop, D. R., Nizkorodov, S. A., Yli-Juuti, T., Schobesberger, S., and Virtanen,
 A.: Insights into the O : C-dependent mechanisms controlling the evaporation of α-pinene
 secondary organic aerosol particles, Atmos. Chem. Phys., 19, 4061-4073, 10.5194/acp-19-4061-2019,
 2019.
- Canagaratna, M., Jayne, J., Jimenez, J., Allan, J., Alfarra, M., Zhang, Q., Onasch, T., Drewnick, F., Coe, H.,
 and Middlebrook, A.: Chemical and microphysical characterization of ambient aerosols with the aerodyne
 aerosol mass spectrometer, Mass Spectrometry Reviews, 26, 185-222, 2007.
- Canonaco, F., Crippa, M., Slowik, J., Baltensperger, U., and Prévôt, A.: SoFi, an IGOR-based interface for the
 efficient use of the generalized multilinear engine (ME-2) for the source apportionment: ME-2 application
 to aerosol mass spectrometer data, Atmospheric Measurement Techniques, 6, 3649, 2013.
- Craven, J. S., Yee, L. D., Ng, N. L., Canagaratna, M. R., Loza, C. L., Schilling, K. A., Yatavelli, R. L. N.,
 Thornton, J. A., Ziemann, P. J., Flagan, R. C., and Seinfeld, J. H.: Analysis of secondary organic aerosol
 formation and aging using positive matrix factorization of high-resolution aerosol mass spectra:





 719
 application to the dodecane low-NO_x system, Atmos. Chem. Phys., 12, 11795-11817,

 720
 10.5194/acp-12-11795-2012, 2012.

721	Cri	ppa, M., Canonaco, F., Lanz, V. A., Äijälä, M., Allan, J. D., Carbone, S., Capes, G., Ceburnis, D., Dall'Osto,
722		M., Day, D. A., DeCarlo, P. F., Ehn, M., Eriksson, A., Freney, E., Hildebrandt Ruiz, L., Hillamo, R.,
723		Jimenez, J. L., Junninen, H., Kiendler-Scharr, A., Kortelainen, A. M., Kulmala, M., Laaksonen, A.,
724		Mensah, A. A., Mohr, C., Nemitz, E., O'Dowd, C., Ovadnevaite, J., Pandis, S. N., Petäjä, T., Poulain, L.,
725		Saarikoski, S., Sellegri, K., Swietlicki, E., Tiitta, P., Worsnop, D. R., Baltensperger, U., and Prévôt, A. S.
726		H.: Organic aerosol components derived from 25 AMS data sets across Europe using a consistent ME-2
727		based source apportionment approach, Atmos. Chem. Phys., 14, 6159-6176, 10.5194/acp-14-6159-2014,
728		2014.
700	T 1	

- Ehn, M., Kleist, E., Junninen, H., Petäjä, T., Lönn, G., Schobesberger, S., Dal Maso, M., Trimborn, A.,
 Kulmala, M., Worsnop, D. R., Wahner, A., Wildt, J., and Mentel, T. F.: Gas phase formation of extremely
 oxidized pinene reaction products in chamber and ambient air, Atmos. Chem. Phys., 12, 5113-5127,
 10.5194/acp-12-5113-2012, 2012.
- Ehn, M., Thornton, J. A., Kleist, E., Sipila, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann,
 R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T., Schobesberger, S.,
 Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurten, T., Nielsen, L. B., Jorgensen, S., Kjaergaard, H.
 G., Canagaratna, M., Dal Maso, M., Berndt, T., Petaja, T., Wahner, A., Kerminen, V.-M., Kulmala, M.,
 Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol,
 Nature, 506, 476-479, 10.1038/nature13032, 2014.
- El Haddad, I., D'Anna, B., Temime-Roussel, B., Nicolas, M., Boreave, A., Favez, O., Voisin, D., Sciare, J.,
 George, C., Jaffrezo, J. L., Wortham, H., and Marchand, N.: Towards a better understanding of the origins,
 chemical composition and aging of oxygenated organic aerosols: case study of a Mediterranean
 industrialized environment, Marseille, Atmos. Chem. Phys., 13, 7875-7894, 10.5194/acp-13-7875-2013,
 2013.
- Fry, J. L., Draper, D. C., Barsanti, K. C., Smith, J. N., Ortega, J., Winkler, P. M., Lawler, M. J., Brown, S. S.,
 Edwards, P. M., Cohen, R. C., and Lee, L.: Secondary Organic Aerosol Formation and Organic Nitrate
 Yield from NO3 Oxidation of Biogenic Hydrocarbons, Environmental Science & Technology, 48, 1194411953, 10.1021/es502204x, 2014.
- Guenther, A., Hewitt, C. N., Erickson, D., Fall, R., Geron, C., Graedel, T., Harley, P., Klinger, L., Lerdau, M.,
 McKay, W. A., Pierce, T., Scholes, B., Steinbrecher, R., Tallamraju, R., Taylor, J., and Zimmerman, P.:
 A GLOBAL-MODEL OF NATURAL VOLATILE ORGANIC-COMPOUND EMISSIONS, Journal of
 Geophysical Research-Atmospheres, 100, 8873-8892, 10.1029/94jd02950, 1995.
- Hakola, H., Tarvainen, V., Bäck, J., Ranta, H., Bonn, B., Rinne, J., and Kulmala, M.: Seasonal variation of
 mono- and sesquiterpene emission rates of Scots pine, Biogeosciences, 3, 93-101, 10.5194/bg-3-93-2006,
 2006.
- Hakola, H., Hellén, H., Hemmilä, M., Rinne, J., and Kulmala, M.: In situ measurements of volatile organic compounds in a boreal forest, Atmos. Chem. Phys., 12, 11665-11678, 10.5194/acp-12-11665-2012, 2012.
- Hari, P., and Kulmala, M.: Station for Measuring Ecosystem–Atmosphere Relations (SMEAR II), Boreal
 Environment Research, 10, 315-322, 2005.
- Huang, S., Rahn, K. A., and Arimoto, R.: Testing and optimizing two factor-analysis techniques on aerosol at Narragansett, Rhode Island, Atmospheric Environment, 33, 2169-2185, <u>https://doi.org/10.1016/S1352-2310(98)00324-0</u>, 1999.
- Huffman, J. A., Docherty, K. S., Aiken, A. C., Cubison, M. J., Ulbrich, I. M., DeCarlo, P. F., Sueper, D., Jayne,
 J. T., Worsnop, D. R., Ziemann, P. J., and Jimenez, J. L.: Chemically-resolved aerosol volatility
 measurements from two megacity field studies, Atmos. Chem. Phys., 9, 7161-7182, 10.5194/acp-9-71612009, 2009.
- Jokinen, T., Sipilä, M., Junninen, H., Ehn, M., Lönn, G., Hakala, J., Petäjä, T., Mauldin Iii, R. L., Kulmala,
 M., and Worsnop, D. R.: Atmospheric sulphuric acid and neutral cluster measurements using CI-APi-TOF,
 Atmospheric Chemistry and Physics, 12, 4117-4125, 10.5194/acp-12-4117-2012, 2012.
- Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M., Yan,
 C., Almeida, J., Troestl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A.,
- 771 Bernhammer, A.-K., Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X., Craven, J., Dias, A., Ehrhart,
- S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen, H.,
 Kangasluoma, J., Kim, J., Krapf, M., Kuerten, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot,





774	S., Molteni, U., Onnela, A., Peraekylae, O., Piel, F., Petaejae, T., Praplan, A. P., Pringle, K., Rap, A.,
775	Richards, N. A. D., Riipinen, I., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Scott, C. E.,
776	Seinfeld, J. H., Sipilae, M., Steiner, G., Stozhkov, Y., Stratmann, F., Tome, A., Virtanen, A., Vogel, A. L.,
777	Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Zhang, X., Hansel,
778	A., Dommen, J., Donahue, N. M., Worsnop, D. R., Baltensperger, U., Kulmala, M., Carslaw, K. S., and
779	Curtius, J.: Ion-induced nucleation of pure biogenic particles, Nature, 533, 521-526, 10.1038/nature17953,
780	2016.
781	Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petäjä, T., Sipilä,
782	M., Schobesberger, S., Rantala, P., Franchin, A., Jokinen, T., Järvinen, E., Äijälä, M., Kangasluoma, J.,
783	Hakala, J., Aalto, P. P., Paasonen, P., Mikkilä, J., Vanhanen, J., Aalto, J., Hakola, H., Makkonen, U.,
784	Ruuskanen, T., Mauldin, R. L., Duplissy, J., Vehkamäki, H., Bäck, J., Kortelainen, A., Riipinen, I., Kurtén,
785	T., Johnston, M. V., Smith, J. N., Ehn, M., Mentel, T. F., Lehtinen, K. E. J., Laaksonen, A., Kerminen, V
786	M., and Worsnop, D. R.: Direct Observations of Atmospheric Aerosol Nucleation, 339, 943-946,
787	10.1126/science.1227385 %J Science, 2013.
788	Lamarque, J. F., Bond, T. C., Evring, V., Granier, C., Heil, A., Klimont, Z., Lee, D., Liousse, C., Mieville, A.,
789	Owen, B., Schultz, M. G., Shindell, D., Smith, S. J., Stehfest, E., Van Aardenne, J., Cooper, O. R.,
790	Kainuma, M., Mahowald, N., McConnell, J. R., Naik, V., Riahi, K., and van Vuuren, D. P.: Historical
791	(1850–2000) gridded anthropogenic and biomass burning emissions of reactive gases and aerosols:
792	methodology and application. Atmos. Chem. Phys., 10, 7017-7039, 10.5194/acp-10-7017-2010, 2010.
793	Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Kurtén, T., Worsnop, D. R., and Thornton, J. A.: An Iodide-
794	Adduct High-Resolution Time-of-Flight Chemical-Ionization Mass Spectrometer: Application to
795	Atmospheric Inorganic and Organic Compounds, Environmental Science & Technology, 48, 6309-6317,
796	10.1021/es500362a, 2014.
797	Lee, B. H., Lopez-Hilfiker, F. D., D'Ambro, E. L., Zhou, P., Boy, M., Petäjä, T., Hao, L., Virtanen, A., and
798	Thornton, J. A.: Semi-volatile and highly oxygenated gaseous and particulate organic compounds observed
799	above a boreal forest canopy, Atmos, Chem. Phys., 18, 11547-11562, 10.5194/acp-18-11547-2018, 2018.
800	Lehtipalo, K., Yan, C., Dada, L., Bianchi, F., Xiao, M., Wagner, R., Stolzenburg, D., Ahonen, L. R., Amorim,
801	A., Baccarini, A., Bauer, P. S., Baumgartner, B., Bergen, A., Bernhammer, AK., Breitenlechner, M.,
802	Brilke, S., Buchholz, A., Mazon, S. B., Chen, D., Chen, X., Dias, A., Dommen, J., Draper, D. C., Duplissy,
803	J., Ehn, M., Finkenzeller, H., Fischer, L., Frege, C., Fuchs, C., Garmash, O., Gordon, H., Hakala, J., He,
804	X., Heikkinen, L., Heinritzi, M., Helm, J. C., Hofbauer, V., Hoyle, C. R., Jokinen, T., Kangasluoma, J.,
805	Kerminen, VM., Kim, C., Kirkby, J., Kontkanen, J., Kürten, A., Lawler, M. J., Mai, H., Mathot, S.,
806	Mauldin, R. L., Molteni, U., Nichman, L., Nie, W., Nieminen, T., Ojdanic, A., Onnela, A., Passananti, M.,
807	Petäjä, T., Piel, F., Pospisilova, V., Quéléver, L. L. J., Rissanen, M. P., Rose, C., Sarnela, N., Schallhart,
808	S., Schuchmann, S., Sengupta, K., Simon, M., Sipilä, M., Tauber, C., Tomé, A., Tröstl, J., Väisänen, O.,
809	Vogel, A. L., Volkamer, R., Wagner, A. C., Wang, M., Weitz, L., Wimmer, D., Ye, P., Ylisirniö, A., Zha,
810	Q., Carslaw, K. S., Curtius, J., Donahue, N. M., Flagan, R. C., Hansel, A., Riipinen, I., Virtanen, A.,
811	Winkler, P. M., Baltensperger, U., Kulmala, M., and Worsnop, D. R.: Multicomponent new particle
812	formation from sulfuric acid, ammonia, and biogenic vapors, 4, eaau5363, 10.1126/sciadv.aau5363 %J
813	Science Advances, 2018.
814	Liebmann, J., Karu, E., Sobanski, N., Schuladen, J., Ehn, M., Schallhart, S., Quéléver, L., Hellen, H., Hakola,
815	H., Hoffmann, T., Williams, J., Fischer, H., Lelieveld, J., and Crowley, J. N.: Direct measurement of NO3
816	radical reactivity in a boreal forest, Atmos. Chem. Phys., 18, 3799-3815, 10.5194/acp-18-3799-2018, 2018.
817	Massoli, P., Stark, H., Canagaratna, M. R., Krechmer, J. E., Xu, L., Ng, N. L., Mauldin, R. L., Yan, C., Kimmel,
818	J., Misztal, P. K., Jimenez, J. L., Jayne, J. T., and Worsnop, D. R.: Ambient Measurements of Highly
819	Oxidized Gas-Phase Molecules during the Southern Oxidant and Aerosol Study (SOAS) 2013, ACS Earth
820	and Space Chemistry, 10.1021/acsearthspacechem.8b00028, 2018.
821	Mohr, C., Lopez-Hilfiker, F. D., Yli-Juuti, T., Heitto, A., Lutz, A., Hallquist, M., D'Ambro, E. L., Rissanen,
822	M. P., Hao, L., Schobesberger, S., Kulmala, M., Mauldin III, R. L., Makkonen, U., Sipilä, M., Petäjä, T.,
823	and Thornton, J. A.: Ambient observations of dimers from terpene oxidation in the gas phase: Implications
824	for new particle formation and growth, 44, 2958-2966, 10.1002/2017gl072718, 2017.
825	Nah, T., Sanchez, J., Boyd, C. M., and Ng, N. L.: Photochemical Aging of α -pinene and β -pinene Secondary
826	Organic Aerosol formed from Nitrate Radical Oxidation, Environmental Science & Technology, 50, 222-
827	231, 10.1021/acs.est.5b04594, 2016.





828	Orlando I I and Tyndall G S. Laboratory studies of organic peroxy radical chemistry: an overview with
829	emphasis on recent issues of atmospheric significance. I Chemical Society Reviews 41, 6294-6317, 2012
830	Pastero P and Tapper II · Positive matrix factorization · A non-negative factor model with ontimal utilization
831	of error estimates of data values Environmetrics 5 111-126 1994
832	Deterior Contracts of data values, Invitation of robust non-negative factor analysis. Chemometrics and Intelligent
833	Laboratory Systems 37 23-35 https://doi.org/10.1016/S0169-7439(96)00044-5 1997
834	Pastero P. The Multilinear Engine A Table Driven Least Squares Program for Solving Multilinear
835	Problems Including the n-Way Parallel Factor Analysis Model Journal of Computational and Graphical
836	Statistics & S42-888 10 1080/10618600 1999 10474853 1999
837	Paciga A. Karnezi F. Kostenidou F. Hildebrandt I. Psichoudaki M. Engelhart G. I. Lee, B. H. Crinna
838	M. Prévôt A. S. H. Baltenspercer II. and Pandis S. N. Volatility of organic aerosol and its components
839	in the megacity of Paris Atmos Chem Phys. 16 2013-2023 10 5194/acn-16-2013-2016 2016
840	Paulson, S. E. and Orlando, J. J.: The reactions of ozone with alkenes: An important source of HOx in the
841	boundary laver. 23, 3727-3730, 10,1029/969103477, 1996.
842	Peräkylä, O., Riva, M., Heikkinen, L., Ouéléver, L., Roldin, P., and Ehn, M.: Experimental investigation into
843	the volatilities of highly oxygenated organic molecules (HOM). Atmospheric Chemistry and Physics
844	Discussions, 2019, 1-28, 10, 5194/acp-2019-620, 2019.
845	Perraud V, Bruns E, A, Ezell M J, Johnson S N, Greaves J, and Finlayson-Pitts B J. Identification of
846	Organic Nitrates in the NO3 Radical Initiated Oxidation of α -Pinene by Atmospheric Pressure Chemical
847	Ionization Mass Spectrometry, Environmental Science & Technology, 44, 5887-5893, 10, 1021/es1005658,
848	2010.
849	Polissar, A. V., Hopke, P. K., Paatero, P., Malm, W. C., and Sisler, J. F.: Atmospheric aerosol over Alaska: 2.
850	Elemental composition and sources, Journal of Geophysical Research: Atmospheres, 103, 19045-19057,
851	1998.
852	Pope III, C. A., Ezzati, M., and Dockery, D. W.: Fine-particulate air pollution and life expectancy in the United
853	States, New England Journal of Medicine, 360, 376-386, 2009.
854	Riva, M., Rantala, P., Krechmer, J. E., Peräkylä, O., Zhang, Y., Heikkinen, L., Garmash, O., Yan, C., Kulmala,
855	M., Worsnop, D., and Ehn, M.: Evaluating the performance of five different chemical ionization
856	techniques for detecting gaseous oxygenated organic species, Atmospheric Measurement Techniques, 12,
857	2403-2421, 10.5194/amt-12-2403-2019, 2019.
858	Sekimoto, K., Koss, A. R., Gilman, J. B., Selimovic, V., Coggon, M. M., Zarzana, K. J., Yuan, B., Lerner, B.
859	M., Brown, S. S., Warneke, C., Yokelson, R. J., Roberts, J. M., and de Gouw, J.: High- and low-
860	temperature pyrolysis profiles describe volatile organic compound emissions from western US wildfire
861	fuels, Atmos. Chem. Phys., 18, 9263-9281, 10.5194/acp-18-9263-2018, 2018.
862	Shiraiwa, M., Ueda, K., Pozzer, A., Lammel, G., Kampf, C. J., Fushimi, A., Enami, S., Arangio, A. M.,
863	Fröhlich-Nowoisky, J., Fujitani, Y., Furuyama, A., Lakey, P. S. J., Lelieveld, J., Lucas, K., Morino, Y.,
864	Pöschl, U., Takahama, S., Takami, A., Tong, H., Weber, B., Yoshino, A., and Sato, K.: Aerosol Health
865	Effects from Molecular to Global Scales, Environmental Science & Technology, 51, 13545-13567,
866	10.1021/acs.est.7b04417, 2017.
867	Song, Y., Shao, M., Liu, Y., Lu, S., Kuster, W., Goldan, P., and Xie, S.: Source apportionment of ambient
868	volatile organic compounds in Beijing, Environmental science & technology, 41, 4348-4353, 2007.
869	Spittler, M., Barnes, I., Bejan, I., Brockmann, K. J., Benter, T., and Wirtz, K.: Reactions of NO3 radicals with
870	limonene and a-pinene: Product and SOA formation, Atmospheric Environment, 40, 116-127,
871	https://doi.org/10.1016/j.atmosenv.2005.09.093, 2006.
872	Stocker, T., Qin, D., Plattner, G., Tignor, M., Allen, S., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley,
813	P:: IPCC, 2013: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to
074 075	the Firth Assessment Report of the Intergovernmental Panel on Climate Change, 1555 pp, in, Cambridge
010 876	Univ. Fless, Cambridge, UK, and New TOFK, 2015. Treast L. Chuang W. K. Cardon H. Hainritzi M. Van C. Maltani H. Ahlm L. Erzas, C. Disseli, F.
010 877	Troesu, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Anim, L., Frege, C., Bianchi, F., Wagner, P., Simon, M. Lahtinala, K. Williamson, C. Cravan, J. S. Dunliagu, J. Adamses, A. Alussida
878	wagner, K., Simon, W., Lenupaio, K., Winnamson, C., Claven, J. S., Dupissy, J., Adamov, A., Almeida, I. Barnhammar, A. K. Braitanlachnar, M. Brilke, S. Diag, A. Ebrhart, S. Elagan, D. C. Eranshir, A.
870	J., Dermannier, AK., Dienemeenne, IVI., Drike, S., Dids, A., Eminart, S., Fidgall, K. C., Franchim, A., Fuchs C. Guida P. Gusal M. Hansal A. Hayla C. D. Jakinan T. Junninan H. Kanasalyama J.
880	ruciis, C., Oulua, K., Oysel, IVI., Hallsel, A., Hoyle, C. K., Jokliell, I., Jullillell, H., Kangasluoma, J., Kackinan H. Kim I. Kranf M. Kuartan A. Laakaanan A. Laular M. Laimingar M. Mathot S.
881	Moehler O Nieminen T Onnela & Petaejae T Dial F M Miattinan D Discanen M D Dondo
882	L. Sarnela N. Schohesberger S. Sengunta K. Sinila M. Smith I. N. Steiner G. Tome Δ. Virtanen
002	2., 541.64, 11, 561.0005001501, 5., 561.54pm, 11, 51pm, 11, 51mm, 5. 11, 510mot, 6., 10mb, A., Viltanon,





A., Wagner, A. C., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Carslaw, K. S., Curtius, J.,
Dommen, J., Kirkby, J., Kulmala, M., Riipinen, I., Worsnop, D. R., Donahue, N. M., and Baltensperger,
U.: The role of low-volatility organic compounds in initial particle growth in the atmosphere, Nature, 533,
527-531, 10.1038/nature18271, 2016.

- Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.: Interpretation of organic components from Positive Matrix Factorization of aerosol mass spectrometric data, Atmos. Chem. Phys., 9, 2891-2918, 10.5194/acp-9-2891-2009, 2009.
- Yan, C., Nie, W., Aijala, M., Rissanen, M. P., Canagaratna, M. R., Massoli, P., Junninen, H., Jokinen, T.,
 Sarnela, N., Hame, S. A. K., Schobesberger, S., Canonaco, F., Yao, L., Prevot, A. S. H., Petaja, T., Kulmala,
 M., Sipila, M., Worsnop, D. R., and Ehn, M.: Source characterization of highly oxidized multifunctional
 compounds in a boreal forest environment using positive matrix factorization, Atmospheric Chemistry and
 Physics, 16, 12715-12731, 10.5194/acp-16-12715-2016, 2016.
- Zha, Q., Yan, C., Junninen, H., Riva, M., Sarnela, N., Aalto, J., Quéléver, L., Schallhart, S., Dada, L.,
 Heikkinen, L., Peräkylä, O., Zou, J., Rose, C., Wang, Y., Mammarella, I., Katul, G., Vesala, T., Worsnop,
 D. R., Kulmala, M., Petäjä, T., Bianchi, F., and Ehn, M.: Vertical characterization of highly oxygenated
 molecules (HOMs) below and above a boreal forest canopy, Atmos. Chem. Phys., 18, 17437-17450,
 10.5194/acp-18-17437-2018, 2018.
- Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Ulbrich, I. M., Ng, N. L., Worsnop, D. R., and Sun, Y.:
 Understanding atmospheric organic aerosols via factor analysis of aerosol mass spectrometry: a review, Analytical and Bioanalytical Chemistry, 401, 3045-3067, 10.1007/s00216-011-5355-y, 2011.
- Zhang, Y., Lin, Y., Cai, J., Liu, Y., Hong, L., Qin, M., Zhao, Y., Ma, J., Wang, X., and Zhu, T.: Atmospheric
 PAHs in North China: spatial distribution and sources, Science of the Total Environment, 565, 994-1000,
 2016.
- Zhang, Y., Cai, J., Wang, S., He, K., and Zheng, M.: Review of receptor-based source apportionment research
 of fine particulate matter and its challenges in China, Science of the Total Environment, 586, 917-929,
 2017.
- Zhang, Y., Peräkylä, O., Yan, C., Heikkinen, L., Äijälä, M., Daellenbach, K. R., Zha, Q., Riva, M., Garmash,
 O., Junninen, H., Paatero, P., Worsnop, D., and Ehn, M.: A novel approach for simple statistical analysis
- 911 of high-resolution mass spectra, Atmospheric Measurement Techniques, 12, 3761-3776, 10.5194/amt-12912 3761-2019, 2019.
- 913